- 1 Phylogenetic and Environmental Context of a Tournaisian Tetrapod Fauna

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25 Summary

26 The end-Devonian to mid-Mississippian time interval has long been known for its 27 depauperate palaeontological record, especially for tetrapods. This interval encapsulates 28 the time of increasing terrestriality among tetrapods, but only two Tournaisian localities 29 previously produced tetrapod fossils. Here we describe five new Tournaisian tetrapods 30 (Perittodus apsconditus, Koilops herma, Ossirarus kierani, Diploradus austiumensis 31 and Avtonerpeton microps) from two localities in their environmental context. A 32 phylogenetic analysis retrieved three taxa as stem tetrapods, interspersed among 33 Devonian and Carboniferous forms, and two as stem amphibians, suggesting a deep split 34 among crown tetrapods. We also illustrate new tetrapod specimens from these and 35 additional localities in the Borders Region of Scotland. The new taxa and specimens 36 suggest that tetrapod diversification was well established by the Tournaisian. 37 Sedimentary evidence indicates that tetrapod fossils are usually associated with sandy 38 siltstones overlying wetland palaeosols. Tetrapods were probably living on vegetated 39 surfaces subsequently flooded. We show that atmospheric oxygen levels were stable 40 across the Devonian/Carboniferous boundary, and did not inhibit the evolution of 41 terrestriality. This wealth of tetrapods from Tournaisian localities highlights the 42 potential for discoveries elsewhere.

43

The term "Romer's Gap" was coined^{1,2} for a hiatus of approximately 25 million years
(Myr)³ in the fossil record of tetrapods from the end-Devonian to the Mid-Mississippian
(Viséan). Following the end-Devonian, the earliest terrestrial tetrapod fauna was known
from the early Brigantian (late Viséan) locality of East Kirkton near Bathgate,

48	Scotland ^{4,5} . By that time, tetrapods were ecologically diverse, and were terrestrially
49	capable. With five or fewer digits, some had gracile limbs ^{6,7} , unlike the polydactylous
50	predominantly aquatic fish-like tetrapods of the Late Devonian ⁸ . Fossils representing
51	transitional morphologies between these disparate forms was almost entirely lacking,
52	limiting both understanding of the acquisition of terrestrial characteristics and the
53	relationships between the diverse mid-Carboniferous taxa. Alternative hypotheses to
54	explain the hiatus have included a low oxygen regime ⁹ or lack of successful collecting
55	in Tournaisian strata ² .

56 Although isolated tetrapod limb bones, girdle elements, and trackways are known from the Tournaisian of the Horton Bluff Formation at Blue Beach, Nova 57 Scotia^{10,11}, only a small fraction has been fully described¹². The only other Tournaisian 58 59 tetrapod material was the articulated skeleton of *Pederpes finneyae*, from the Tournaisian Ballagan Formation near Dumbarton, western Scotland^{13,14}. More recently, 60 new taxa from this formation in the Borders Region of Scotland were reported², but 61 further collecting from five localities (Supplementary Fig. 1) has since produced more 62 63 data about the fauna, its environment, and climatic conditions. 64 Our analysis shows that the Tournaisian included a rich and diverse assemblage 65 of taxa which included close relatives of some Devonian forms on the tetrapod stem, and basal members of the amphibian stem. We diagnose, name and analyse five taxa (Figs 1-66 5), and summarize at least seven others that are distinct but undiagnosable at present 67 68 (Fig. 6, Supplementary Figs 2-6).

69 Tetrapods occupied a juxtaposed mosaic of microhabitats including ponds,
70 swamps, streams, and floodplains with highly variable salinity and water levels in a

71	sharply contrasting seasonal climate. Their fossils are most closely associated with
72	palaeosols and the overlying sandy siltstones. These indicate exposed and vegetated land
73	surfaces that were then flooded ^{15,16} (Supplementary Fig 7). This varied environment
74	persisted over the 12 million years of the Tournaisian ³ . We show that atmospheric
75	oxygen levels were stable across the Devonian/Carboniferous boundary, and did not
76	therefore compromise terrestrial faunal life (contra ref 9).
77	Differential diagnoses below give the characters in which each differs from all
78	other tetrapods in its combination of autapomorphic and derived (relative to Devonian
79	taxa) characters.
80	This published work and the nomenclatural act it contains have been registered in
81	Zoobank: http://www.zoobank.org:pub:4BFFB544-7B0B-4F2F-80EC-11226C0FDAAB
82	Tetrapoda Goodrich, 1930 indet.
83	Perittodus apsconditus gen. et sp. nov. Clack and Smithson T.R. Fig. 1 e-g.
84	Smithson et al., 2012 (fig. 4), new taxon A.
85	LSID. urn:lsid:zoobank.org:act 69DB72E5-F9BD-49C6-B471-CD8E03767732
86	Etymology. Genus from perittos (Greek) 'odd' and odus (Greek) 'tooth' referring to the
87	unusual dentition of the mandible. Species from apsconditus (Latin) 'covert, disguised,
88	hidden, secret or concealed', referring to the fact that key parts were only discovered by
89	micro-CT scanning.
90	Holotype. UMZC 2011.7.2 a and b. Cheek region of skull, lower jaw, and postcranial
91	elements in part and counterpart.
92	Locality and Horizon. Willie's Hole, Whiteadder Water near Chirnside. Ballagan
93	Formation. Early mid Tournaisian.
94	Diagnosis. Autapomorphies: unique adsymphysial and coronoid dentition –
95	adsymphysial with two tusks and at least two smaller teeth, anterior coronoid with two

96 or three larger tusks, middle coronoid with two larger and two or three smaller teeth,

97 posterior coronoid row of small teeth; lozenge-shaped dorsal scales bearing concentric

98 ridges centred close to one edge nearer to one end. Derived characters: deeply excavated

99 jugal with narrow suborbital bar; lateral line an open groove on jugal.

100 Plesiomorphies & characters of uncertain polarity: No mesial lamina of postspenial

101 (state of angular not known); 35 dentary teeth including spaces; 29 maxillary teeth

102 including spaces; room for possibly 6 teeth on premaxilla; marinal teeth similar in size;

103 short broad phalanges, rounded unguals longer than wide with ventral ridge.

104 Attributed specimen. UMZC 2016.1. Isolated dentary and adsymphysial (in micro-CT

scan) from Burnmouth Ross end cliffs, 373.95 m above the base of the Ballagan

106 Formation. Mid Tournaisian.

107 **Remarks**: Lower jaw length 68 mm. Maxilla of holotype visible in micro-CT scan.

108 UMZC 2016.1 is almost identical in size and dentition to the holotype. The pattern is

109 most similar to but not identical with, that of the Devonian taxon *Ymeria*¹⁷. A distinct

110 denticulated ridge on the prearticular is set off from the remainder of the bone by a

111 ventral groove. Radius and ulna are of approximately equal length. A partial ischium

- 112 reveals similarities to that of $Baphetes^{18}$.
- 113

114 *Koilops herma* gen. et sp. nov. Clack and Smithson T.R. Fig. 1 a-b.

115 Smithson et al., 2012 (fig. 2C), 'probable new taxon'.

116 LSID. urn:lsid:zoobank.org:act 8C43E66A-3822-49B4-B3B5-E43C79FA9C70

117 Etymology. Genus from *koilos* (Greek) 'hollow or empty', and *ops* (Greek) 'face',

118 referring to the skull mainly preserved as natural mould. Species from *herma* (Greek)

119 'boundary marker, cairn, pile of stones'. The specimen, from the Borders Region of

120 Scotland, has transitional morphology between Devonian and Carboniferous tetrapods.

121 Holotype. NMS G. 2013.39/14. Isolated skull mainly as a natural mould.

122 Locality and Horizon. Willie's Hole, Whiteadder Water near Chirnside. Ballagan

123 Formation. Early mid Tournaisian.

124 **Diagnosis.** Autapomorphies: fine irregular dermal ornament with conspicuous curved

125 ridges around the parietal foramen and larger pustular ornament anterior to parietal

126 foramen. Derived characters: deeply excavated jugal with narrow suborbital bar; large

127 parietal foramen.

128 Plesiomorphies & characters of uncertain polarity: Orbit oval with slight anterior

129 embayment; prefrontal-postfrontal contact narrow, anterior to orbit mid-length; about 8

130 premaxillary teeth recurved, sharply pointed, ridged towards base; closed palate,

131 denticulated pterygoid; vomers bearing tusks and smaller teeth, at least four moderately

132 large teeth on palatine; short rounded snout, only slightly longer than maximum orbit

133 length.

134 Remarks. Skull length 80 mm. The dermal bones are robust and well integrated so the135 individual was almost certainly not a juvenile.

136

137 Ossirarus kierani gen. et sp. nov. Clack and Smithson T.R. Fig. 2.

138 LSID. urn:lsid:zoobank.org:act FC9FAB5C-CC3E-4D0D-B7D7-8030FBAA4F0C

139 Etymology. Genus from ossi (Latin) 'bones' and rarus (Latin) 'scattered or rare.'

140 Specific name to honour Oliver and Betty Kieran, representing the Burnmouth

141 community, who have supported us and encouraged local interest and co-operation.

142 Holotype. UMZC 2016.3. A single block containing scattered skull and postcranial

143 remains.

144 Locality and Horizon. Burnmouth Ross end cliffs, 340.5 m above the base of the

145 Ballagan Formation. Mid Tournaisian.

146 **Diagnosis.** Autapomorphies: tabular elongate triangle forming a conspicuous tabular

147 horn with a convex lateral margin. Derived character: tabular-parietal contact;

148 exoccipital separate from basioccipital.

149 Plesiomorphies & characters of uncertain polarity: Jugal with extensive posterior

150 component, with anteriorly placed shallow contribution to orbit; lozenge-shaped

151 interclavicle; humerus with elongate and oblique pectoralis process comparable with the

152 ventral humeral ridge of elpistostegalians and *Acanthostega*; multipartite vertebrae with

153 diplospondylous widely notochordal centra and neural arches as unfused bilateral

halves.

155 **Remarks:** Estimated skull length 50 mm based on comparisons with *Acanthostega*,

156 *Ichthyostega* and *Greererpeton*¹⁹⁻²¹. The primitive jugal morphology, with an elongated

157 postorbital region and an anteriorly placed orbital margin contributing less than 25% of

158 the orbit margin, is similar to that in *Acanthostega*¹⁹ and *Ichthyostega*²⁰. The tabular has

an elongated posterior process, but its lateral margin does not show an embayment for a

160 spiracular notch. The bones are robust, with well defined overlap areas for

161 interdigitating sutures. Though disarticulated, these suggest that the individual was not a

162 juvenile. The specimen shows the earliest known occurrence of a separate exoccipital.

163

164 *Diploradus austiumensis* gen. et sp. nov. Clack and Smithson T.R. Fig. 3.

165 *LSID*. urn:lsid:zoobank.org:act 268DDD4F-289D-4F83-8172-1A18A1007B7C

166 Etymology. Genus from *diplo* (Greek) 'double' and *radus* (Greek) 'row' referring to the

167 double coronoid tooth row. Species from *austium* (Latin) 'mouth of a river or stream'

168 referring to Burnmouth.

169 Holotype. UMZC 2015.55.4. Small disrupted skull with lower jaw, palate and skull

170 roofing bones.

171 Locality and Horizon. Burnmouth Ross end cliffs, 373.95 m above the base of the 172 Ballagan Formation. Mid Tournaisian.

Diagnosis. Autapomorphies: lower jaw with irregular double row of denticles along the coronoids; around 51 dentary teeth and spaces, with enlarged tusk at position 3 and the 174 175 largest teeth in positions 8-13; parietals short, pineal foramen anteriorly placed; ?narrow 176 curved pre- and postfrontals. Derived characters: deeply excavated jugal with narrow 177 suborbital bar; parasphenoid with broad, flattened posterior portion with lateral wings, 178 earliest known occurrence of a parasphenoid crossing the ventral cranial fissure, 179 cultriform process flat, narrow. 180 Attributed specimen. UMZC 2016.4 a and b. The anterior end of a mandible from 341 181 m above the base of the Ballagan formation at Burnmouth. 182 Plesiomorphies & characters of uncertain polarity: Unsutured junction between 183 prearticular and splenial series; adductor fossa dorsally placed; adsymphysial plate 184 possibly lacking dentition; closed, denticulated palate; broad pterygoid, quadrate ramus

185 narrow with vertically orientated medial ascending lamina; ossified hypotranchial

186 elements; maxilla and premaxilla with spaces for 35 and 10-12 teeth respectively;

187 maxilla-premaxilla contact narrow, lacking interdigitations; dermal ornament with low

188 profile, irregular on skull table, ridged on squamosal and quadratojugal.

Remarks. Lower jaw length 30 mm, superficially resembling that of Sigournea²², 189

- 190 although a relationship is not supported by cladistic analysis. The thinness of the bones
- 191 and their distribution suggest a juvenile.

192

173

193 Avtonerpeton microps gen. et sp. nov. Otoo, Clack and Smithson T.R. Fig. 4.

194 LSID. urn:lsid:zoobank.org:act E1E094A8-FAC0-4A2A-A13D-487D7775FBE1

195 Etymology. Genus name from Ayton, the parish in the Scottish Borders from which the

196 specimen came, and *erpeton* (Greek) 'crawler' or 'creeping one'. Species name from

197 *micro* (Greek) 'small' and *ops* (Greek) 'face'.

198 Holotype. UMZC 2015.55.8. Partial skull and scattered postcrania visible only in micro-

199 CT scan (Supplementary Movie Files)

Locality and Horizon. Burnmouth Ross end shore exposure, 340.6 m above the base of
the Ballagan Formation. Mid Tournaisian.

202 **Diagnosis.** Autapomorphies: two enlarged premaxillary teeth plus one large tooth space

at posterior end of premaxilla; 5 teeth on premaxilla; adsymphysial with a single tooth;

204 coronoids apparently lacking shagreen; L-shaped lacrimal; vomer with at least one tooth,

205 palatine with one large fang but lacking smaller teeth; ectopterygoid with at least two

206 teeth and possible smaller teeth. Derived characters shared with colosteids: course of

207 lateral line on maxilla and nasal; dentary teeth larger and fewer than upper marginal

208 teeth; single large Meckelian fenestra; interpterygoid vacuities longer than wide; single

209 large parasymphysial fang on dentary; ilium with a single strap-shaped iliac process.

210 **Remarks**. Reconstructed skull length about 50 mm. Other distinguishing features: short

snout, approximately similar in length to orbit diameter; naris and choana both very

212 large relative to skull size – relatively larger than in *Greererpeton*. The enlarged

213 premaxillary teeth prefigure those of more derived colosteids^{e.g.21}, but the dentary lacks

the corresponding reciprocal notch. This appears an early expression of a feature that

215 becomes more elaborate in later taxa. All coronoids bear at least one tooth. Some

216 colosteids lack coronoid teeth, and instead bear shagreen, a variable condition among

217 individuals²³. The small size of the skull but the strong integration of the lower jaw

218 bones suggest a subadult or adult in which case the large orbit is unlikely to be a

219	juvenile feature (c.f. juvenile <i>Greererpeton</i> CMNH 11095 ²⁴). Its gracile limbs,
220	metapodial bones and phalanges resemble Colosteus rather than Greererpeton.
221	Clavicular ornament is similar to that of other colosteids ^{25,26} . The single iliac process is
222	shared with other colosteids and with temnospondyls. The earliest known occurrence of
223	this feature.
224	
225	Results
226	Cladistic Analysis
227	We performed parsimony and Bayesian analyses of a new data matrix (Supplementary
228	Data Character list and Data matrix) incorporating the new tetrapods. No taxon could
229	be safely deleted ²⁷ . Parsimony with all characters unordered and equally weighted
230	produced 4718 shortest trees, a poorly resolved strict consensus (Fig. 5, Supplementary
231	Fig. 8), and moderate branch support.
232	Four parsimony analyses with implied weighting, each using a different value (3,
233	4, 5, 10) of the concavity constant K^{28} produced many fewer trees (Fig. 5a, b), with
234	novel topologies and increased stability for most of the new taxa. In these analyses, the
235	relative positions of Ossirarus, Perittodus, and Diploradus remain unaltered (Methods
236	and Supplementary Fig. 8). Except in the analysis with K=10, Koilops and
237	Aytonerpeton emerge as stem amphibians ^{29-31, but see 32,33} with Aytonerpeton close to
238	Tulerpeton+colosteids. With characters reweighted by their rescaled consistency index,
239	all new taxa emerge as stem tetrapods.
240	We also performed a Bayesian analysis (Fig. 5c). The results were largely similar
241	to the parsimony analysis, except for the position of Ossirarus. In the Bayesian analysis,
242	Ossirarus appears as a stem amniote, whilst Perittodus, Diploradus, Koilops, and

243 *Aytonerpeton* are stem tetrapods.

Despite inconsistencies, these results imply a substantial reshuffling of the branching sequence of Carboniferous stem tetrapods relative to previous studies²⁹⁻³³, with interspersed Carboniferous and Devonian taxa pointing to a more ramified stem of tetrapod diversification. If corroborated by further evidence, a firmer placement of *Aytonerpeton* and *Koilops* within crown tetrapods would suggest a deep split between stem amphibians and stem amniotes within the Tournaisian..

251 Geology and Environment

The Ballagan Formation (Inverclyde Group) underlies much of the Midland Valley ofScotland and the northern margin of the Northumberland Basin. At Burnmouth the

vertically dipping strata probably span the entire Tournaisian^{2,34}. Environmental

interpretation was based on a 490 m core from a borehole through the formation, a

complete logged succession at centimetre scale intervals through 520 m at Burnmouth,

and an 8 m section at Willie's Hole (Fig. 6, Methods and Supplementary Fig. 7).

258 *Perittodus apsconditus* occurs within a 6 cm thick laminated grey siltstone¹⁶ that 259 contains a network of cracks filled with sandy siltstone identical to that of the overlying 260 bed. Occurring within laminated siltstones, this may record an autochthonous lake 261 dweller. Associated fossils comprise plants, actinopterygians, myriapods and ostracods. 262 *Koilops* occurs within a unit comprising four beds of alternating black and green 263 siltstone in which abundant palaeosol clasts indicate erosion and transport of land-264 surface sediment during flooding events.

265 Diploradus occurs in a 40 cm thick, bedded, black sandy siltstone that lies 266 between pedogenically modified grey siltstones. Associated fossils comprise fish scales, 267 abundant plant fragments, megaspores, and shrimp and scorpion cuticle. 268 Ossirarus and Aytonerpeton occur within a complex 15 cm thick grey-black 269 sandy siltstone that overlies a gleyed palaeosol and grades upwards into a laminated 270 grey siltstone with brecciation cracks (Fig. 6, Methods and Supplementary Fig. 7). 271 Ossirarus occurred just above the palaeosol in a light grey clay-rich sandy siltstone, 272 whereas Avtonerpeton occurred within an overlying black sandy siltstone with abundant 273 plant material. Associated fauna comprise abundant plants, megaspores, unusually 274 abundant rhizodont bones and scales, actinopterygians, chondrichthyans (Ageleodus, 275 Gyracanthus), dipnoans, eurypterids and ostracods. 276 An association between wetland palaeosols and tetrapod-bearing facies has 277 emerged from our studies, significant because those horizons indicate a vegetated land surface (Fig. 6) 15,16 . The flood-plain environments of semi-permanent water bodies, 278 279 marsh, river banks and areas of dry land with trees were laid down at a time of change in 280 the land plant flora of the Mississippian following the end-Devonian extinctions. The new flora initiated a change in fluvial and floodplain architecture³⁵⁻³⁷. Progymnosperms 281

had been almost eliminated in the extinctions, but thickets and forests were re-

established in the early-mid Tournaisian with lycopods as the dominant flora. At

284 Burnmouth many beds with abundant spores of the creeping lycopod Oxroadia include

tetrapods. Terrestrial ground-dwelling arthropods, such as myriapods and scorpions

286 fossils of which have been found at Burnmouth and at Willie's Hole, formed a possible

food supply for tetrapods..

288

289 Atmospheric oxygen levels in the Tournaisian

To address the low oxygen hypothesis⁹ we examined fossil charcoal (fusinite) in the
Ballagan Formation to compare atmospheric oxygen levels in the Tournaisian with the
Late Devonian and later Mississippian.

293 Charcoal, either as microscopic dispersed organic matter (DOM) or visible in 294 hand specimens is relatively common at Burnmouth and Willie's Hole. Although 295 charcoal is reported from the Tournaisian Horton Bluff Formation, Nova Scotia³⁸ as 296 indicating O_2 concentrations above 16%, no quantitative study to validate this result has 297 been undertaken.

298 We analysed DOM from 73 rock samples from Burnmouth shore and Willie's 299 Hole. For comparison with wildfire activity before and after Romer's Gap, we also 300 analysed 42 samples from the Viséan of East Fife, Scotland (Strathclyde Group) and 9 301 samples from the Famennian of Greenland (Stensiö Bjerg Formation) (Supplementary 302 Fig. 9 and Supplementary Table 1). All were found to contain fusinite, with a mean 303 abundance relative to total phytoclasts of 2.2%, 2.3% and 2.6% for the Famennian, 304 Tournaisian and Viséan, respectively. We also analysed 12 samples from Willie's Hole 305 which had a mean value of 2.0% (Supplementary Table 1). Not only do these results 306 mean that fire activity persisted through Romer's Gap and indicate that atmospheric O₂ 307 did not fall below 16%, but also that there was no significant change in charcoal 308 production compared with the Famennian and Viséan (Supplementary Fig. 9). This 309 strongly suggests that atmospheric O₂ was stable across this time interval, directly refuting hypoxia⁹ as an explanation for Romer's Gap. 310

311

312 Discussion

Although an extinction event at the end of the Devonian saw the demise of many archaic
fish groups³⁹, our studies provide new perspectives on the recovery and diversification
of surviving groups, which went on to found the basis of modern vertebrate
diversity^{40,41}.

The new tetrapods show no close relationship to each other, exhibiting different combinations of plesiomorphic and derived characters. Some taxa cluster with Devonian forms, suggesting a possible relict fauna, whereas others appear more crownward, even clustering near the base of the crown group. They imply an early radiation of tetrapods during the Tournaisian, and at the same time, suggest a blurring of the Devonian-Carboniferous (D-C) boundary in respect of tetrapod evolution, a feature also noted in tetrapod remains from Nova Scotia¹².

If confirmed, our results imply a deep split between stem amphibians and stem 324 325 amniotes in the earliest Carboniferous. This accords with most molecular dates for the split that place it at an average of 355 Ma 42,43 a date only 4 Ma after the end-Devonian. 326 327 It suggests that the origin of the tetrapod crown group occurred soon after the extinction 328 event as tetrapods began to recover. Their radiation into a range of new taxa parallels that of lungfish⁴⁰ and chondrichthyans⁴¹ as they adapted to a post-extinction world. 329 The occurrence of probable plesiomorphic members of the Crassigyrinidae² and 330 331 Colosteidae indicates an inception 20-24 Myr earlier than the Late Mississippian as previously considered⁴⁴. Other tetrapod material of uncertain attribution are distinct and 332

increase known tetrapod diversity in the Tournaisian (Fig. 6 and Supplementary Figs 2-6).

335 The preponderance of small animals throughout the sequence is unusual, notably 336 a very small tetrapod in a horizon 33 m above the D-C boundary, around 1 Myr after the 337 extinction event (Fig. 6). None of the five taxa described above has a skull length of 338 more than 80 mm. This could indicate preservational or collector bias, but they occur 339 throughout different lithologies, horizons and localities (Fig. 6 and Supplementary Figs 340 2-6). Larger tetrapod taxa are found at Willie's Hole, about one quarter of the way up 341 the sequence, probably representing about 3 or 4 Myr above the D-C boundary. Larger 342 sizes seem to have appeared relatively rapidly in the Tournaisian, as also documented by trackways³⁸ and challenge suggestions of a prolonged period of reduced body size in 343

344 vertebrates following the DC extinction event 45 .

The tetrapods of the Ballagan Formation lived in a mosaic of floodplain environments. Some were under water for long periods, others alternated between land surface and standing water. A recent study of the development of *Polypterus* shows how early in life, their skeletons can be differentially modified in response to exposure to water-based or land-based conditions⁴⁶. Such skeletal flexibility might have contributed to the origin of tetrapod terrestrial morphology in the varied environments of the Ballagan Formation.

The wealth and diversity of tetrapod taxa from the Tournaisian refutes the proposal of depauperate Tournaisian stage, and our charcoal studies show that atmospheric oxygen levels, stable from the Famennian to the Viséan, were not a causal factor for the apparent gap. We emphasise the importance of exploring or re-exploring non-marine Tournaisian sites elsewhere in the world, and examining previously overlooked lithologies.

358

359 References

360	1.	Coates, M. I & Clack, J. A. Romer's Gap – tetrapod origins and terrestriality.
361		Bull. Mus. Nat. Hist. Nat. 17, 373-388 (1995)
362	2.	Smithson, T. R., Wood, S. P., Marshall, J. E. A. & Clack, J. A. Earliest
363		Carboniferous tetrapod and arthropod faunas from Scotland populate Romer's
364		Gap. Proc. Natl. Acad. Sci. USA 109, 4532-4537 (2012)
365	3.	Cohen, K.M., Finney, S.C., Gibbard, P.L. & Fan, JX. The International
366		Chronostratigraphical Chart, Episodes 36, 199-204 (2013)
367		http://stratigraphy.org/ICSchart/ChronostratChart2016-04.pdf
368	4.	Wood, S. P., Panchen, A. L. & Smithson, T. R. A terrestrial fauna from the
369		Scottish Lower Carboniferous. Nature 314, 355-356 (1985)
370	5.	Rolfe, W. D. I., Clarkson, E. N. K. & Panchen, A. L. (Eds). Volcanism and early
371		terrestrial biotas. Trans. R. Soc. Edinb. Earth Sci. 84, (1994)
372	6.	Milner, A. R. & Sequeira, S. E. K. The temnospondyl amphibians from the
373		Viséan of East Kirkton, West Lothian, Scotland. Trans. R. Soc. Edinb. Earth Sci.
374		84, 331-362 (1994)
375	7.	Smithson, T. R., Carroll, R. L., Panchen, A. L. & Andrews, S. M. Westlothiana
376		lizziae from the Viséan of East Kirkton, West Lothian, Scotland. Trans. R. Soc.
377		Edinb. Earth Sci. 84, 417-431 (1994)
378	8.	Clack, J. A. Gaining Ground: The origin and evolution of tetrapods. 2 nd Ed. 1-
379		523. (Indiana Univ. Press, 2012)
380	9.	Ward, P. D., Labandeira, C., Laurin, M. & Berner, R. A. Confirmation of
381		Romer's Gap as a low oxygen interval constraining the timing of initial
382		arthropod and vertebrate terrestrialisation. Proc. Natl. Acad. Sci. USA 103,
383		16818-16822 (2006)

384	10. Carroll, R. L., Belt, E. S., Dineley, D. L., Baird, D. & McGregor, D. C.Excursion
385	A59, Vertebrate palaeontology of Eastern Canada. 24 th International Geological
386	Congress, Montreal (1972)
387	11. Clack, J. A. & Carroll, R. L. in Amphibian Biology, Vol. 4: Palaeontolog (eds
388	Heatwole, H. & Carroll, R. L.) 1030-1043 (Surrey Beatty, 2000)
389	12. Anderson, J. S., Smithson, T. R., Mansky, C. F., Meyer, T. & Clack, J. A. A
390	diverse tetrapod fauna at the base of Romer's Gap. Plos One
391	DOI:10.1371/journal.pone.0125446 (2015)
392	13. Clack, J. A. An early tetrapod from 'Romer's Gap'. Nature 418, 72-76 (2002)
393	14. Clack, J. A. & Finney, S. M. Pederpes finneyae, an articulated tetrapod from the
394	Tournaisian of western Scotland. J. Syst. Palaeont. 2, 311-346 (2005)
395	15. Kearsey, T. I., Bennett, C. E., Millward, D., Davies, S. J., Gowing, C. J. B.,
396	Kemp, S. J., Leng, M. L., Marshall, J. E. A., Browne, M. A. E. The terrestrial
397	landscapes of tetrapod evolution in earliest Carboniferous seasonal wetlands of
398	SE Scotland. Palaeogeogr, Palaeoclim, Palaeoeco 457, 52-69
399	doi.org/10.1016/j.palaeo.2016.05.033 (2016)
400	16. Bennett, C. E., Kearsey, T. I., Davies, S. J., Millward, D. Clack, J. A.,
401	Smithson, T. R. & Marshall, J. E. A. Early Mississippian sandy siltstones
402	preserve rare vertebrate fossils in seasonal flooding episodes. Sediment.
403	"Enhanced Article"; doi: 10.1111/sed.12280 (2016)
404	17. Clack, J.A., Ahlberg P.E., Blom H., & Finney S.M. A new genus of Devonian
405	tetrapod from East Greenland, with new information on the lower jaw of
406	Ichthyostega. Palaeont. 55, 73-86 (2012)

407	18. Milner, A. C. & Lindsay, W. Postcranial remains of Baphetes and their bearing
408	on the relationships of the Baphetidae (= Loxommatidae). Zoo. J. Linn. Soc. 122,
409	211-235 (1998)
410	19. Clack, J. A. The dermal skull roof of Acanthostega, an early tetrapod from the
411	Late Devonian. Trans. R. Soc. Edinb. Earth Sci. 93, 17-33 (2002)
412	20. Jarvik, E. The Devonian tetrapod Ichthyostega. Fossils and Strata 40, 1-206
413	(1996)
414	21. Smithson, T. R. The cranial morphology of Greererpeton burkemorani
415	(Amphibia: Temnospondyli). Zoo. J. Linn. Soc. 76, 29-90 (1982)
416	22. Lombard, R. E. & Bolt, J. R. Sigournea multidentata, a new stem tetrapod from
417	the Upper Mississippian of Iowa, USA. J. Palaeont. 80, 717-725 (2006)
418	23. Bolt J. R. & Lombard, R. E. The mandible of the primitive tetrapod
419	Greererpeton, and the early evolution of the tetrapod lower jaw. J. Paleont. 75,
420	1016-1042 (2001)
421	24. Godfrey, S. J. Ontogenetic changes in the skull of the Carboniferous tetrapod
422	Greererpeton burkemorani Romer 1969. Phil. Trans. Roy. Soc. Lond. B 323,
423	135-153 (1989a)
424	25. Hook, R. W. Colosteus scutellatus (Newberry) a primitive temnospondyl
425	amphibian from the Middle Pennsylvanian of Linton, Ohio. Am. Mus. Novit.
426	2770, 1-41 (1983)
427	26. Godfrey, S. J. The postcranial skeletal anatomy of the Carboniferous tetrapod
428	Greererpeton burkemorani Romer 1969. Phil. Trans. Roy. Soc. Lond. B 323, 75-
429	133 (1989b)
430	27. Wilkinson, M. Coping with abundant missing entries in phylogenetic inference
431	using parsimony. Syst. Biol. 44, 501-514 (1995)

- 432 28. Goloboff, P. A. Estimating character weights during tree search. *Cladistics* 9, 83433 89 (1993)
- 434 29. Ruta, M., Coates, M. I. & Quicke, D. L. J. Early tetrapod relationships revisited.
 435 *Biol. Rev.* 78, 251-345 (2003)
- 30. Ruta, M. & Clack, J. A. A review of *Silvanerpeton miripedes*, a stem amniote
 from the Lower Carboniferous of East Kirkton, West Lothian, Scotland. *Trans. R. Soc. Edinb. Earth Sci.* 97, 31-63 (2006)
- 439 31. Klembara J., Clack, J. A. & Milner A. R. Cranial anatomy, ontogeny, and
- relationships of the Late Carboniferous tetrapod *Gephyrostegus bohemicus*Jaekel, 1902. *J. Vert. Pal.* 34, 774-792. (2014).
- 442 32. Laurin, M. The evolution of body size, Cope's rule and the origin of amniotes.
 443 *Syst. Biol.* 53, 594-622 (2004).
- 444 33. Marjanović, D. & Laurin, M. The origin(s) of extant amphibians: a review with
- 445 emphasis on the "lepospondyl hypothesis". *Geodiversitas* **35**, 207-272. (2013)
- 446 34. Greig, D. C. *Geology of the Eyemouth district*. Memoir of the British Geological
- 447 Survey, Sheet 34. (1988)
- 448 35. Garcia, W.J., Storrs, G.W. & Greb, S.F. The Hancock County tetrapod locality:
- 449 A new Mississippian (Chesterian) wetlands fauna from western Kentucky
- 450 (USA). Geol. Soc. Am. Spec. Papers **399**, 155-167 (2006)
- 451 36. Davies, N. S. & Gibling, M. R. The sedimentary record of Carboniferous rivers:
- 452 Continuing influence of land plant evolution on alluvial processes and
- 453 Palaeozoic ecosystems. *Earth-Sci. Rev.* **120**, 40-79 (2013)
- 454 37. Corenblit, D., Davies, N. S., Steiger, J., Gibling, M. R. & Bornette, G.
- 455 Considering river structure and stability in the light of evolution: feedbacks
- 456 between riparian vegetation and hydrogeomorphology. *Earth Surf. Proc. Land.*

- **457 40**, 189-207 (2015)
- 458 38. Mansky, C. F. & Lucas, S. G. Romer's Gap revisited: continental assemblages
 459 and ichno-assemblages from the basal Carboniferous of Blue Beach, Nova
 460 Scotia, Canada. *Bull. New. Mex. Mus. Nat. Hist.* 60, 244-273 (2013).
- 39. Sallan, L. C. & Coates, M. I. End-Devonian extinction and a bottleneck in the
 early evolution of modern jawed vertebrates. *Proc. Nat. Acad. Sci. USA* 107,
 10131-10135 (2010)
- 464 40. Smithson, T. R., Richards, K. R. & Clack, J. A. Lungfish diversity in Romer's
 465 Gap: reaction to the end-Devonian extinction. *Palaeont.* 59, 29-44 (2016).
- 466 41. Richards, K. R., Sherwin, J. E., Smithson, T. R., Bennion, R. F., Davies, S. J.,
- 467 Marshall, J. E. A. & Clack, J. A. A new fauna of early Carboniferous
- 468 chondrichthyans from the Scottish Borders. <u>http://www.palass.org/meetings-</u>
- 469 <u>events/annual-meeting/2015/annual-meeting-2015-cardiff-poster-abstracts</u>
 470 (2015)
- 471 42. Hedges, S. B., Marin, J., Suleski, M., Paymer, M. & Kumar, S. Tree of Life
 472 reveals clock-like speciation and diversification. *Mol. Biol. Evol.* 32, 835-845
 473 (2015)
- 474 43. Kumar S. & Hedges S. B. TimeTree2: species divergence times on the iPhone.
 475 *Bioinformatics* 27:2023-2024 www.timetree.org (2011)
- 476 44. Clack, J. A., Witzmann, F., Snyder D., and Müller, J. 2012 A colosteid-like early
 477 tetrapod from the St. Louis Limestone (Early Carboniferous, Meramecian), St.
 478 Louis, Missouri, USA. *Fieldiana Life Earth Sci.*5, 17-39.
- 479 45. Sallan L. C. & Gallimberti, A. K. Body-size reduction in vertebrates following
 480 the end-Devonian mass extinction. *Science* 350, 812-815 (2015)

481	46. Standen, E. M., Du, T. Y. & Larsson, C. E. Developmental plasticity and the
482	origin of tetrapods. Nature 513, 54-58. (2014) doi:10.1038/nature13708
483	47. Ahlberg, P. E. and Clack, J. A. Lower jaws, lower tetrapods: a review based on
484	the Devonian tetrapod Acanthostega. Trans. R. Soc. Edinb. 89, 11-46 (1998)
485	48. Clack, J. A. The Scottish Carboniferous tetrapod Crassigyrinus scoticus
486	(Lydekker) - cranial anatomy and relationships. Trans. R. Soc. Edinb. 88, 127-
487	142. (1998)
488	49. Goloboff, P. A., Farris, J. S., & Nixon, K. C. TNT, a free program for
489	phylogenetic analysis. Cladistics 24, 1-13 (2008)
490	50. Carpenter, D. K., Falcon-Lang, H. J., Benton, M. J., & Henderson, E.
491	Carboniferous (Tournaisian) fish assemblages from the Isle of Bute, Scotland:
492	systematics and palaeoecology. Palaeont. 57, 1215-1240 (2014)
493	51. Friedman, M. & Sallan, L. C. Five hundred million years of extinction and
494	recovery: a Phanerozoic survey of large-scale diversity patterns in fishes.
495	Palaeont. 55, 707-742 (2012)
496	52. Andrews, J. E., Turner, M. S., Nabi, G. & Spiro, B. The anatomy of an early
497	Dinantian terraced floodplain: palaeo-environment and early diagenesis.
498	Sediment. 38, 271-287 (1991)
499	53. Turner, M.S. Geochemistry and diagenesis of basal Carboniferous dolostones
500	from southern Scotland (Unpublished Ph.D. thesis, University of East Anglia,
501	1991)
502	54. Belt, E. S., Freshney, E. C. & Read, W. A. Sedimentology of Carboniferous
503	cementstone facies, British Isles and Eastern Canada. J. Geol. 75, 711-721.
504	(1967)
505	55. Scott, W. B. The sedimentology of the cementstone group in the Tweed basin:

- 506 Burnmouth and the Merse of Berwick. (Unpublished PhD Thesis, Sunderland 507 Polytechnic 1971) 508 56. Scott, W. B. Nodular carbonates in the Lower Carboniferous, Cementstone 509 Group of the Tweed Embayment, Berwickshire: evidence for a former sulphate 510 evaporite facies. Scot. J. Geol. 22, 325-345 (1986) 511 57. Barnet, A. J., Wright, V. P. & Crowley, S. F. Recognition and significance of 512 paludal dolomites: Late Mississippian, Kentucky, USA. Interl Assn Sediment. 513 Spec. Publ. 45, 477-500 (2012) 514 58. Muchez, P. & Viaene, W. Dolocretes from the Lower Carboniferous of the 515 Campine-Brabant Basin, Belgium. Pedologie 37, 187–202 (1987) 516 59. Searl, A. Pedogenic dolomites from the Oolite Group (Lower 517 Carboniferous), South Wales. Geol. J. 23, 157–169 (1988) 518 60. Vanstone, S. D. Early Carboniferous (Mississippian) palaeosols from southwest 519 Britain: influence of climatic change on soil development. J. Sediment. Petrol. 520 **61**, 445–457 (1991) 521 61. Wright, V. P., Vanstone, S. D. & Marshall, J. D. Contrasting flooding histories 522 of Mississippian carbonate platforms revealed by marine alteration effects in 523 palaeosols. Sedimentology 44, 825–842 (1997) 524 62. Wood, G., Gabriel, A.M. & Lawson, J.C. Palynological techniques -processing 525 and microscopy. 29–50 in: Palynology: Principles and Applications, Volume 1. 526 Principles. (eds) Jansonius, J. and McGregor, D. C. (American Association of 527 Stratigraphic Palynologists Foundation, Texas, 1996). 528 63. American Society for Testing and Materials (ASTM). D2799 - 13. Standard test 529 method for microscopical determination of the maceral composition of coal. in:
- 530 Annual book of ASTM standards section 5 Petroleum products, lubricants, and

531	their fossil fuels. Volume 05.06 Gaseous Fuels; Coal and Coke; Bioenergy and
532	Industrial Chemicals from Biomass. (West Conshohocken, PA, ASTM
533	International. DOI: 10.1520/D2799-13
534	http://www.astm.org/Standards/D2799.htm (2013)
535	64. Hansen, K. W. & Wallmann, K. Cretaceous and Cenozoic evolution of seawater
536	composition, atmospheric O ₂ and CO ₂ : A model perspective. Am. J. Sci. 303, 94-
537	148 (2003).
538	65. Bergman, N. M., Lenton, T. M. & Watson, A. J. COPSE: a new model of
539	biogeochemical cycling over Phanerozoic time. Am. J. Sci. 304, 397-437 (2004).
540	66. Arvidson, R.S., Mackenzie, F.T. & Guidry, M. Magic: A Phanerozoic model for
541	the geochemical cycling of major rock-forming components. Am. J. Sci. 306,
542	135–190 (2006).
543	67. Berner, R. A. GEOCARBSULF: A combined model for Phanerozoic
544	atmospheric O ₂ and CO ₂ . Geochim. Cosmochim. Ac. 70, 5653–5664 (2006)
545	68. Berner, R. A. Phanerozoic atmospheric oxygen: new results using the
546	GEOCARBSULF model. Am. J. Sci. 309, 603-606 (2009)
547	69. Robinson, J. M. Phanerozoic atmospheric reconstructions: a terrestrial
548	perspective. Palaeogeogr. Palaeoclim. Palaeoecol. 97, 51-62 (1991)
549	70. Scott, A. J. & Glasspool, I. J. The diversification of Paleozoic fire systems and
550	fluctuations in atmospheric oxygen concentration. Proc. Nat. Acad. Sci. USA
551	103, 10861-10865 (2006)
552	71. Glasspool, I. J., & Scott, A. C. Phanerozoic concentrations of atmospheric
553	oxygen reconstructed from sedimentary charcoal. Nature Geoscience. 3, 627-630
554	(2010)

555	72. Glasspool, I. J., Scott, A. C., Waltham, D., Pronina, N. & Shao, L. The impact of
556	fire on the Late Paleozoic Earth system. Front. Plant Sci. 6, 1-13 (2015)
557	73. Belcher, C. M., Yearsley, J. M., Hadden, R. M., McElwain, J. C. & Guillermo,
558	R. Baseline intrinsic flammability of Earth's ecosystems estimated from
559	paleoatmospheric oxygen over the past 350 million years. Proc. Nat. Acad. Sci.
560	<i>USA</i> 107 , 22448-22453 (2010)
561	74. Tyson, R. V. Sedimentary organic matter. 1-615. Chapman & Hall, London,
562	(1995)
563	75. Scott, A. J. & Glasspool, I. J. Observations and experiments on the origin and
564	formation of the inertinite group macerals. Internl J. Coal Petr., 70, 53-66.
565	(2007)
566	76. Owens, B., McLean, D. and Simpson, K. R. M. Reappraisals of the
567	Mississippian palynostratigraphy of the East Fife coast, Scotland, United
568	Kingdom. Palynology, 29, 23–47 (2005)
569	77. Forsythe, J. H. and Chisholme, J. I. The geology of East Fife (explanations of the
570	Fife portion of 'one-inch' geological sheet 41 and part of sheet 49). (Her
571	Majesty's Stationary Office, Edinburgh, 284pp. 1977)
572	78. Stephenson, M. H., Williams, M., Monghan, A. A., Arkley, S., Smith, R. A.
573	Dean, M., Browne, M. A. E. and Leng, M. Palynomorph and ostracod
574	biostratigraphy of the Ballagan Formation, Midland Valley of Scotland, and
575	elucidation of intra-Dinantian unconformities. Proc. Yorks. Geol. Soc. 55, 131-
576	143 (2004)
577	79. Marshall, J. A. E., Astin, T. R. and Clack, J. A. East Greenland tetrapods are
578	Devonian in age. Geology, 27, 637–640 (1999)

579	80. Astin, T. R., Marshall, J. E. A., Blom, H. and Berry, C. M. The sedimentary
580	environment of the Late Devonian East Greenland tetrapods. 93 – 109 in: Vecoli,
581	M., Clement, G. and Meyer-Berthaud, B. (eds.) The terrestrialization process:
582	modelling complex interactions at the biosphere-geosphere interface. Geol. Soc.
583	Lond. Spec. Publ. 339 (2010)
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587	Figure legends
588	Figure 1. a-b Koilops herma gen. et sp. nov. (National Museum of Scotland NMS G.
589	2013.39/14). a , Photograph of specimen, mainly preserved as natural mould. b ,
590	Interpretive drawing of specimen. c-g, Perritodus apsconditus gen. et sp. nov.
591	(University Museum of Zoology, Cambridge UMZC 2011.7.2a). c, Photograph of main
592	specimen block. d, Reconstruction of lower jaw in external view, from scan data and
593	part and counterpart specimens. e, Reconstruction of lower jaw in internal view made
594	from scan data and part and counterpart specimens. f, Segmented model from scans of
595	lower jaw in internal view. g, segmented model from scans of lower jaw in internal
596	view. Colour code in f,: orange, dentary; red, adsymphysial plate; turquoise, part of
597	prearticular; yellow, first coronoid; blue, second coronoid; cerise, third coronoid; pink,
598	splenial; violet, angular; purple, prearticular; green, splenial; external bones greyed out.
599	In g, green, splenial. Scale bar in a, b, and c, 10 mm. Abbreviations: add foss, adductor
600	fossa; adsymph, adsymphysial; ang, angular; cor, coronoid; dent, dentary; ecto,
601	ectopterygoid; fro, frontal; intemp, intertemporal; jug, jugal; l, left; lac, lacrimal; llc,
602	lateral line canal; max, maxilla; oa, overlap area for pterygoid; pal, palatine; par,
603	parietal; pofr, postfrontal; porb, postorbital; pospl, postsplenial; preart, prearticular;

604 prefro, prefrontal; premax, premaxilla; psph, parasphenoid; pteryg, pterygoid; quad,

605 quadrate; quj, quadratojugal; surang, surangular; vom, vomer.

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617

618 Figure 3. Diploradus austiumensis gen. et sp. nov. (UMZC 2015.55.4). a, Photograph of

619 complete specimen. Scale bar 10 mm, **b**, Map of specimen showing distribution of

620 elements, c, Drawing of right maxilla, d, Upper, interpretive drawing of specimen;

621 lower, reconstruction of jaw in internal view. e, Drawing of parasphenoid. f, Drawing of

622 right jugal in internal view. **g**, Drawing of skull table. **h**, Drawing of pterygoid in dorsal

623 view. Scale bars in **b-h**, 5 mm. Abbreviations as for Figures 1 and 2 except for: nat

- 624 mould popar, natural mould of postparietal.
- 625

626 Figure 4. Aytonerpeton microps gen. et sp. nov. (UMZC 2015.55.8). a, Still from micro-

627 CT scan of block containing most of the specimen. **b**, Interpretive drawing of right side

628 of skull and palate. c, Stills from micro-CT scan of right lower jaw in (upper image)

629 dorsal view and (lower image) mesial view. d, Still from micro-CT scan of right palate 630 in approximately ventral view. e, Still from micro-CT scan of entire specimen in the 631 main block. Arrows point to elements in g. f, Enlargement of ilium in lateral (left image) 632 and medial (right image) views. g, Elements of hind limb. In c, and d, note the sutures 633 between pterygoid and marginal palatal bones, and the lower jaw bones, are tightly 634 sutured and difficult to see in the scan. Abbreviations as for Figures 1 and 2, except for: 635 mar Meck fen, margin of Meckelian fenestra; sym, symphysis; septomax, septomaxilla. 636 Scale bars for all except **f** are 10 mm. Scale bar for f is 5 mm. 637

638 Figure 5. Three cladograms: two from TNT analysis and one from Bayesian analysis. a,

639 Single most parsimonious tree obtained from implied weights search with k=3 (see text

and Supplementary Data for details). **b**, strict consensus of four equally parsimonious

641 trees obtained from implied weights search with k=4. **c**, Bayesian analysis tree. See main

642 text, methods, and Supplementary Data for details.

643

644 Figure 6. Burmmouth sedimentary log showing palaeosol and tetrapod fossil

645 distribution. Left hand column shows the sedimentary log for Burnmouth with the

646 tetrapod horizons indicated. Right hand column shows the distribution of palaeosols and

647 their thicknesses. Photographs a-g show some of the tetrapod specimens found in

648 addition to those in Figs 1-4.

649 Specimen **a**, an isolated jugal (UMZC 2016.13) from the same bed that yielded the

650 partial *Crassigyrinus*-like jaw in ref 2, horizon approximately 383 m above the base of

- the Ballagan Formation. This is a thick localized conglomerate lag containing many
- isolated vertebrate bones, plant remains and charcoal. The shape of the jugal is unique
- among the tetrapods so far collected from the Ballagan, in its relative contribution to the

654 orbit margin. Probable new taxon 1. Specimens **b-f**, tetrapod specimens from a closely 655 juxtaposed set of horizons beyond the resolution of the log to differentiate, between 340-656 341m above the base of the Ballagan: **b**, an isolated tetrapod maxilla (UMZC 2016.9); **c**, 657 tetrapod belly scales (UMZC 2016.12) and metapodials/phalanges (UMZC 2016.10, 11); 658 d, skull bones and belly scales (UMZC 2016.8); e, Micro-CT scan of the two 659 overlapping bones in **d**. They are probable frontal bones of a *Pederpes*-like tetrapod; **f**, 660 partial skull table and postorbitals from slightly above the Burnmouth horizon yielding 661 Avtonerpeton microps (UMZC 2016.7). Probable new taxon 2? May be associated with 662 those in Supplementary Fig. 2, but not with Avtonerpeton. Scale bar 10 mm. (Micro-663 CT by K. Z. Smithson); g, phalanges or metapodials and skull elements of a small 664 tetrapod from Burnmouth (UMZC 2016.5 a, b). Probable new taxon 3. Left hand image, 665 largest elements circled. Right hand image, dentigerous bone near top left corner. Other 666 elements include a probable jugal and rib fragments (not figured). These remains are the 667 earliest post-Devonian tetrapod specimens found in the UK. They come from a horizon 668 approximately 33m above the base of the Ballagan Formation that was probably 669 deposited about 1 Myr after the start of the Carboniferous. Scale bars for all except g are 670 10 mm. Scale for for g is 5 mm. (Photographs by J. A. Clack) 671

672 METHODS

673 Micro-CT data

674 Specimen UMZC 2016.3 Ossirarus and NMS G. 2013.39/14 Koilops and UMZC

675 2011.7.2a *Perittodus* were prepared mechanically with mounted needle, some matrix

676 was removed from Ossirarus with a brush and water, consolidated where necessary with

677 Paraloid B72. Specimens UMZC 2011.7.2a Perittodus and UMZC 2015.55.8

678 *Aytonerpeton* were scanned at the Cambridge Tomography Centre with a Nikon

- 679 XTH225 ST scanner. Scan data:- *Perittodus*: Isotropic voxel size, 0.0444mm.
- 680 Projections:1080, Filter: 0.25mm Cu, Xray kV:160, Xray μA: 70, Slices:1647.
- 681 Exposure: 1000, Gain: 24 dB. UMZC 2015.55.8 Aytonerpeton: Isotropic voxel size:
- 682 0.0609mm. Projections: 1080, Filter: None, Xray kV: 120, Xray μA: 125, Slices: 1789,
- 683 Exposure: 1000, Gain: 24 dB. .
- 684

685 Cladistic analysis

A new database of 46 taxa coded for 214 osteological characters (170 cranial, 43

687 postcranial), and was subjected to maximum parsimony analyses. It was designed to

688 include representative early tetrapods. Characters were drawn up to capture the features

of the new taxa as far as possible in the context of the range of early tetrapods available

690 for comparison. Most were drawn from recent analyses 14,29-31,44,47,48. Some

691 characters were reworded or reformulated and all were independently scored by JAC

from personal observation or from the literature. These were checked for accuracy by

693 MR. Characters are arranged in alphabetical order grouped into regions of the anatomy

694 (Supplementary Data Character list and Data Matrix).

The data matrix was subjected to maximum parsimony analyses in TNT v. 1.1^{49} . 695 696 Several experiments of taxon and character manipulation were carried out, as detailed 697 below, with identical search protocols throughout. Given the size of the matrix, tree 698 searches relied on heuristic algorithms, following a simple series of steps under the 699 'Traditional search' option in the 'Analyze' menu in TNT. Before each search, we 700 modified memory requirements under the 'Memory' option in the 'Settings' menu. One 701 hundred Mbytes of general RAM were allocated, and a total of 50,000 trees were 702 selected as the maximum size of tree space for the exploration of alternative tree 703 topologies. In the initial part of the 'Traditional search' ('Wagner trees' box ticked), we

chose 10,000 replicates (random stepwise addition sequences of taxa), keeping a
maximum of five trees at the end of each replicate, using the bisection-reconnection
algorithm for tree branch swapping, and retaining all trees found at the end of all
replicates. A new round of branch swapping was then applied to all trees retained from
the initial search ('trees from RAM' box ticked). For each set of experiments, where
applicable, we summarized the results in the form of a strict consensus, a 50% majorityrule consensus.

711 Using the search settings expounded above, we carried out three types of 712 parsimony analysis. The first parsimony analysis, employing all taxa and characters 713 from the original matrix, treated all characters as having equal unit weight (default TNT 714 option). The second analysis, again using all taxa and characters, was based on implied character reweighting²⁸, briefly described as follows. Given a character, its implied 715 716 weight (W) is given by K / (K + M - O), where M and O represent, respectively, the 717 greatest number of character-state changes and the observed number of character-state 718 changes for that character. The constant of concavity (K) is an integer, the value of 719 which determines the most parsimonious trees as those trees for which W is maximized 720 across all characters. As the selection of K is arbitrary, we experimented with increasing 721 values (K = 3, 4, 5 and 10) (Fig 5, Supplementary Fig. 8). We did not report details of 722 searches with other K values, as our goal was to establish whether the Tournaisian taxa 723 showed stable positions within a minimal range of implied weighting increments. 724 However, we ran analyses with values varying between 6 and 10, with mixed outcomes. 725 In some cases, the Tournaisian taxa are heavily reshuffled, in others the branching 726 sequence of other groups revealed implausible arrangements that, we feel, were dictated 727 by varying amounts of homoplasy in the data, although a proper characterization of this 728 phenomenon requires further testing. Topologies with K=10 are reported as an example.

In the third analysis, characters were reweighted by the maximum value (best fit)
of their rescaled consistency indexes, such as were obtained from the first analysis.

Statistical branch support was evaluated through character resampling via
bootstrap (resampling with replacement; ref.) and jackknife (resampling without
replacement, with 33% of characters removed; ref.), using 1000 replicates in each case
and collapsing nodes with less than 50% support.

Of all the new Tournaisian taxa, only *Diploradus* appears in a maximum agreement subtree (a taxonomically pruned tree showing only taxa for which all most parsimonious trees agree upon relationships).

As for the implied weighting analysis, we found stable mutual arrangements for most Tournaisian taxa with K = 3, 4 and 5. With K = 10, the branching sequence of Tournaisian taxa differed from those found with smaller values. In addition, slightly different branching patterns emerge for various early tetrapod taxa/groups following different implied weighting searches. Below, we highlight key differences among various tree topologies.

744 In trees generated with K = 3, 4 and 5, Ossirarus, Perittodus and Diploradus 745 emerge as increasingly crownward taxa, in that sequence, along the tetrapod stem group, 746 whilst Aytonerpeton and Koilops are placed among stem amphibians and are thus part of 747 the tetrapod crown group. Ossirarus is crownward of a (Ventastega + Ichthyostega) 748 clade, with Ossinodus placed either immediately anti-crownward of (K = 3), in a 749 polytomy with (K = 4), or immediately crownward of Ossirarus (K = 5). Perittodus is 750 the sister taxon to the Devonian Ichthyostega-like taxon Ymeria, and the (Perittodus + 751 Ymeria) clade forms the sister group to Pederpes. Diploradus is immediately crownward of a (Whatcheeria + Occidens) clade, which in turn occurs crownward of (Pederpes + 752 753 (Perittodus + Ymeria)). However, the branching sequence of Carboniferous stem

754 tetrapods more crownward than *Diploradus* varies. Thus, in trees with K = 3, the branching sequence includes Crassigvrinus, Doragnathus, (Megalocephalus + 755 756 *Baphetes*) and *Loxomma*. In trees with K = 4, the sequence includes only *Crassigyrinus* 757 and *Doragnathus*, whereas all baphetids form a clade on the amphibian stem 758 (Megalocephalus + (Loxomma + Baphetes)). In trees with K = 5, the baphetid clade is, 759 once again, on the amphibian stem, but the sequence of stem tetrapods crownward of 760 Diploradus differs substantially, and includes (Eucritta + Doragnathus), Sigournea and 761 *Crassigvrinus*. In trees from K = 3 and 4, the (*Aytonerpeton + Sigournea*) clade forms 762 the sister group to a (*Koilops* + (*Tulerpeton* + (*Greererpeton* + *Colosteus*))) clade. In 763 turn, this wider group joins temnospondyls on the amphibian stem, with Caerorhachis as 764 a more immediate sister taxon. In trees from K = 5, Aytonerpeton is collapsed in a 765 trichotomy with temnospondyls and the (Koilops + (Tulerpeton + (Greererpeton + 766 *Colosteus*))) clade. With K = 10, the results match those from the second set of 767 parsimony analyses (reweighting). 768 As for other tetrapod groups, the amniote stem undergoes little reshuffling in 769 trees derived from different K values. The most noticeable difference among such trees 770 is the placement of *Silvanerpeton* and *Gephyrostegus*, both of which are immediately 771 crownward of the 'anthracosauroids' (*Eoherpeton* + (*Pholiderpeton* + *Proterogyrinus*)) 772 but swap their positions as the first and second most crownward plesion after 773 anthracosauroids. 774 With characters reweighted by the maximum value of the rescaled consistency 775 index, we found three trees differing only in the relative positions of Whatcheeria, 776 Pederpes and Occidens, all of which form a clade. In those trees, all new Tournaisian

- taxa appear on the tetrapod stem. In particular, *Aytonerpeton* and *Perittodus* are sister
- taxa, and together they join *Ymeria*. In crownward order, the sequence of stem tetrapods

includes: Acanthostega, Ossinodus, Ventastega, Ichthyostega, Ossirarus, the (Ymeria

780 (Aytonerpeton + Perittodus)) clade, the (Whatcheeria, Pederpes, Occidens) clade,

781 Diploradus, Doragnathus, Sigournea, a (Koilops + (Tulerpeton + (Greererpeton +

782 *Colosteus*))) clade, *Crassigyrinus*, and a baphetid clade. *Caerorhachis* and *Eucritta*

- appear as the earliest diverging plesions on the amphibian and amniote stem groups,
- 784 respectively.
- 785

786 Sedimentological and Environmental Interpretation

787 The borehole was located at Norham near Berwick-Upon-Tweed (British National Grid

Reference [BNGR] 391589, 648135), and the Burnmouth section is at BNGR 396000-

789 <u>661000</u>.

790 The stratigraphical position of the succession at Willie's Hole is inferred from a 791 nearby borehole (Hutton Hall Barns, BGS Registered number NT85SE1. The exact 792 stratigraphical position of the Willie's Hole (WH) section is uncertain within the overall 793 succession. No direct correlation with the succession recorded in the Hutton Hall Barns 794 borehole is possible because the borehole is an old one and the level of detail 795 insufficient, plus the fact that distinctive markers are not present in the Ballagan 796 Formation. However, that borehole proved 142.5m of Ballagan Formation strata - the log 797 is good enough to define precisely where the base is, resting on Kinnesswood Formation. 798 The proximity of WH to the borehole allows us to infer that the WH section lies 799 approximately 150m above the base of the Ballagan Formation. The palynological 800 samples from WH contained Umbonatisporites distinctus, a spore that is only found in 801 the lower part of our borehole core. We argue that therefore the WH section belongs to 802 the lower part of the Ballagan Formation. We indicated some uncertainty in the figure 803 and gave an approximate range.

804 The dominance of actinopterygians and rhizodonts within these lakes indicates brackish-freshwater salinity levels ^{50,51}. Diverse palaeosols ¹⁵ and palynology suggest 805 806 habitats including forest, low-growing and creeping flora, wetland and desiccating pools 807 traversed by rivers (predominantly meandering channels) and saline-hypersaline lakes depositing cementstones and evaporites (Fig. 6 and Supplementary Fig. 7)^{27-31,52} The 808 809 saline-hypersaline lake deposits in the Ballagan Formation have been interpreted to represent brackish marginal marine or hypersaline⁵²⁻⁵⁶ conditions. Other dolomitic units 810 from the Mississippian are interpreted as saline coastal marshes⁵⁶⁻⁶¹. Erosive-based, 811 812 cross-bedded sandstone units (one to tens of metres thick) with basal conglomerate lags cut into all other facies 34 . The lags contain disarticulated vertebrate material including 813 acanthodian, rhizodont and tetrapod bones¹⁶. 814

815

816 Charcoal Analysis

Dispersed organic matter (DOM) was extracted by standard palynological
demineralisation techniques⁶². Measurement of maceral reflectance in oil was by means
of a Zeiss UMSP 50 Microspectrophotometer, housed in the School of Ocean and Earth
Science, National Oceanography Centre Southampton, University of Southampton
Waterfront Campus. Measurements were made under standard conditions as defined by
the International Committee for Coal Petrology⁶³.

Model-based estimates of atmospheric oxygen concentration during the early Tournaisian vary from 10 - 20%, with more recent models favouring the higher figure⁶⁴⁻⁶⁸. As an alternative, fossil charcoal (fusinite) is used by several authors as a

826	proxy for atmospheric oxygen $^{69-72}$, as wildfire activity, and hence charcoal production,
827	is proportional to oxygen supply ⁷³ . Controlled burning experiments ⁷³ have
828	demonstrated that when O_2 exceeds the present atmospheric level (PAL) of 20.9%, fire
829	activity rapidly increases and reaches a plateau at around 24%; therefore, we infer that
830	fusinite abundance is likely insensitive to any further increase. Conversely, fire activity
831	is strongly supressed below 20% O_2 and switched off completely below 16%, even in
832	very dry conditions ⁷³ . The most comprehensive attempt thus far to reconstruct
833	Phanerozoic O_2 in this way ⁶⁹ indicated 25.6% O_2 during Romer's Gap – substantially
834	higher than PAL and exceeding the presumed upper limit of fusain sensitivity (24%).
835	However, this study was based on the inertinite (= microscopic fusinite) content of
836	coals, which are infrequent during the Tournaisian, so sampling density was relatively
837	low. Furthermore, we assume that large-scale forest fires will have a far greater
838	influence on coal deposits, formed in situ in forest mires, than on the more distal
839	deposits of the kind examined here.
840	By focusing on DOM extracted from sedimentary rocks other than coal, fusinite
841	content can be measured through stratigraphic successions in which coals are rare or
842	absent. The values reported here represent the proportion of fusinite within the organic
843	matter isolated from each 5g shale sample, based on examination of 500 organic (i.e.

plant derived) macerals. This indicates the proportion of plant-derived material in the
sample which has been burned at high temperatures, and is therefore independent of
sediment supply.

847 The specific Famennian and Viséan sampling localities chosen were selected because, as848 well as being of the required age:

849

• The stratigraphic context of the sampled formations is well understood, with

850	well-established biozonation (Supplementary Table 1).
851	• Thermal maturity in these successions is low. This is essential, because with
852	increasing thermal maturity the reflectance of non-pyrolitic macerals (most
853	notably vitrinite) increases, eventually rendering them indistinguishable from
854	fusinite.
855	• Both localities represent largely terrestrial environments, containing a succession
856	of fluviodeltaic, lacustrine or nearshore marine deposits (Supplementary Table
857	1). Sediments deposited in such environments represent an accumulation point
858	for river-transported organic material derived from the wider region; this
859	mitigates the distorting effect of local fire activity,
860	The organic maceral fusinite is considered synonymous with charcoal and can be
861	distinguished from other maceral types by its reflectance under incident light ⁷⁴ ; we have
862	focused solely on fusinite for this study because, although most other members (semi-
863	fusinite) of the inertinite group are also accepted as pyrolitic in origin ⁷⁵ , their
864	reflectance forms a continuum between that of vitrinite and fusinite and forms the bulk
865	of the organic matter. This makes the % sum of semi-fusinite and fusinite very large
866	(>90%) and less reliable.
867	Supplementary Data Table 1b gives the samples taken from Famennian sites,
868	Burnmouth, Willie's Hole and Visean sites. These were analysed for charcoal content.
869	Mean abundance was 2.0%, which is within error of data obtained from Burnmouth
870	Shore, suggesting that the contribution from local fire activity (if any) was similar at
871	both sites (Supplementary Table 1 and Supplementary Fig. 9).
872	
873	Data availabilitity statement. Correspondence and requests for materials should be

874 addressed to Jennifer A. Clack j.a.clack@zoo.cam.ac.uk. Specimen information is
available from the respective housing institutions. Micro-CT scan data to be placed inthe NERC National Geoscience Data Centre.

877

878 Author contributions

879 JAC is corresponding author and lead PI and with TRS, JAC, BKAO, and KZS

880 collected, described and analysed the tetrapod specimens. CEB, TIK, SJD and DM

881 contributed to the stratigraphical, sedimentological and environmental studies. JEAM,

882 DKC, and EJR contributed to the charcoal, palynological and stratigraphical studies. MR

and JAC contributed to the phylogenetic analysis. AJR contributed information on the

884 arthropods, SAW provided additional work on micro-CT scan data. AJR, SAW and

885 NCF organised the Willie's Hole excavation that provided sedimentological

886 information. All authors contributed to discussion, preparation and writing the paper.

887

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37

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905











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Supplementary Information contains

Supplementary Video file 1 (.mov). This shows the skull of *Aytonerpeton* as a rotational view of the 3-D skull in the main text Figure 4. (12.7 MB)

Supplementary Video file 2 (.mov). This shows the whole block containing *Aytonerpeton* including the skull and postcranial elements as a rotational view in 3-D.

Supplementary Figures

Fig. 1. Locality map.

Figs 2-6. Figures of additional tetrapod specimens. The tetrapod fossils in this collection represent conservatively a sample of at least 7 new taxa, but possibly more. A further taxon is represented by NMS G.2012.39.22 ("Ribbo" in ref 2). The *Crassigyrinus*-like jaw UMZC 2011.9.1 in ref 2 may belong to one of those figured here, or to NMS G.2012.39.22, although the dermal ornamentation does not match that in any of them. Fig. 7. Detailed log of the most productive section of the Burnmouth sequence, and the changing palaeoenvironment that it represents.

Fig. 8. Three additional cladograms.

Fig. 9. Box plot of fusinite abundances from the Famennian – Visean.

Supplementary Data. This contains a pdf of the character list with the sources for the characters and the data matrix used in the cladistic analysis.

Supplementary Table shows fusinite sampling sites, numbers of samples and environments, and fusinite abundances from the Famennian, Tournaisian and Visean.

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Supplementary Figure 1: Map showing distribution of tetrapod-bearing localities in Scotland. Inset – Borders Region on the east coast where most of the finds have been made.



Supplementary Figure 2: Associated skull and other bones (UMZC 2016. 6 a-d) from the same Burnmouth horizon yielding *Aytonerpeton microps* (main text Fig.4). Probable new taxon 2.

a, Micro-CT scan of bones including a parasphenoid in ventral view, a dentary or maxilla, a jugal, a probable postorbital, a possible clavicle, centrum and ribs. **b**, The parasphenoid in dorsal view showing the dorsum sellae. **c**, Photograph of the external surface of the block with a close-up of the parasphenoid and jaw elements. The dentition and jugal bones are unlike those of *Aytonerpeton*, and the parasphenoid is unlike those of either colosteids or temnospondyls from later in the Carboniferous. The short dorsum sellae is similar to that of some later temnospondyls, but the extent of the denticulation on the posterior plate of the parasphenoid is unique. (Photograph by T. R. Smithson, micro-CT by K. Z. Smithson



Supplementary Figure 3: Two skulls representing new taxa from Willie's Hole (SPW bed 2).

a, A small skull (NMS G. 2012.39.77) in part and counterpart. Probable new taxon 3. The skull is highly fractured, each part and counterpart showing both palatal and skull roofing bones. The position of the orbits cannot be ascertained. The drawing in **b**, is based on information from both part and counterpart. Notable is the evidence for a closed palate in a skull with a strongly embayed squamosal for a spiracular or possibly otic notch, a unique combination of features in a Tournaisian tetrapod. There is almost no overlap in the preserved bones with those of *Ossirarus*, but it is clear that it does not share the distinctive tabular with that taxon, and its dermal ornament is quite different. Although NMS G. 2012.39.77 is smaller than *Ossirarus*, it is equally well ossified. 10 mm scale bar on photographs. (Photographs by T. R. Smithson); **c**, Part of a skull roof in natural mould, from Willie's Hole (NMS G. 2012.39.95). Probable new taxon 4. The specimen has associated skull bones, phalanges and an ilium. The shape of the parietal foramen, the impression left by its surrounding bones, and its proportions distinguish it from *Koilops herma*. Scale bar 10 mm. (Photograph by J. A. Clack)



Supplementary Figure 4: UMZC 2011.7.16, a large tetrapod in part and counterpart from Willie's Hole (SPW bed 2). Probable new taxon 5. The specimen includes an interclavicle similar to that of *Pederpes*, and fragments of the skull, numerous ribs and belly scales, a few centra, and epipodials, none of which resembles those of *Pederpes*. The ribs do not bear flanges, the centra are less well ossified, and the epipodials are much more slender. There are also rectangular and lightly ornamented probable dorsal scales (inset, arrows. Inset also shows belly scales). This specimen does not resemble that of NMS G. 2012.39.22 ("Ribbo" in ref 2) in which the ribs are longer and more curved. Scale bar in inset, 10 mm. (Photographs by J. A. Clack and T. R. Smithson)



Supplementary Figure 5: UMZC 2011.7.13, a large jaw and scattered skull bones from Willie's Hole (SPW bed 2). Probable new taxon 5?

a, Isolated skull bones collected from the specimen by S. P. Wood include a tabular that is similar to that of *Crassigyrinus*, although there are no other obvious similarities between the two. **b**, Whole specimen. **c**, Belly scales, similar to those in UMZC 2011. 7.16. **d**, Teeth show a recurved tip and lateral keels that are not found in other tetrapods from the Tournaisian collection. However, there are no teeth preserved in UMZC 2011. 7.16, so that attribution of these two specimens to the same taxon cannot be ruled out. Scale bars in **a-c** are 10 mm, in **d**, 5 mm. (Photographs by J. A. Clack and T. R. Smithson)



Supplementary Figure 6. a,b, A small humerus from Tantallon Castle, Gin Head (NMS G.2016.15.1). Probable new taxon 7. **c,** A maxilla in part and counterpart from the Heads of Ayr.

a, views left to right: ventral; posteroventral; anterior. **b**, from a micro-CT scan, views left to right: dorsal, ventral, posterior, anterior. Note the large deltopectoral process (dpc), placed anterior to a recognizable shaft, and the extent of torsion between proximal and distal ends. The conspicuous latissimus dorsi process (ldp) is shared with *Baphetes* and *Pederpes*, but the overall shape of the humerus is quite different from those two. The matrix is a crudely bedded coarse volcaniclastic sedimentary rock, containing ostracods, bivalves and fish remains. **c**, Maxilla from Heads of Ayr, possibly similar to that in main text Fig. 6. Scale bars 10 mm (Photographs by T. R. Smithson, micro-CT images by K. Z. Smithson)



Supplementary Figure 7: Sedimentary conditions associated with the tetrapods. Palaeoenvironment of two of the tetrapod deposits. Left: Sedimentary log of partial section at Burnmouth with *Aytonerpeton* and *Ossirarus*, from 332 to 356 metres above the base of the Ballagan Formation. Between the sandstone units at the top and base of this section the sedimentary rocks comprise an overbank facies association. This succession records the transition from wet to dry conditions through time, with environments illustrated in the reconstructions for dry and wet periods (right). The tetrapod fossil-bearing horizons within this section are sandy siltstones.



Supplementary Figure 8: Three cladograms. **a**, Majority rule consensus of 4718 trees, unweighted analysis; **b**, Strict consensus of 4 trees, implied weights analysis with k=5; **c**, Strict consensus of 4 trees, implied weights analysis with k=10. The majority rule consensus is of identical topology to this, and the simple weighted analysis almost identical.



Supplementary Figure 9. Visean fusinite abundance in Euramerica, based on analysis of material from the Stensiö Bjerg Formation (Greenland), Burnmouth Shore/Willie's Hole, and the Strathclyde Group (Fife, Scotland). White bars indicate mean values.



Character List

Unless otherwise stated, characters drawn from publications below, with their numbers

*Ahlberg & Clack 1998⁷⁴; \triangle Clack 1998⁷⁵; \$Ruta et al. 2002²⁹; #Clack & Finney 2004¹⁴; %Ruta & Clack, 2006³⁰; @Klembara et al., 2014³¹; \diamond Clack et al. 2012⁴⁴ Skull roof and braincase

- Anterior tectal (accessory dermal bone associated with naris having surface ornament and absent lateral line canal; treated here as septomaxilla): present = 0, absent = 1 \$8; #1
- 2. Anterior tectal: narial opening ventral to it = 0: narial opening anterior to it = 1 # 2
- Basioccipital: indistinguishable from exoccipitals = 0, separated by suture = 1 #3;
 %242
- 4. Basioccipital: ventrally exposed portion longer than wide = 0, shorter than wide = 1
 #4; %243
- 5. Basioccipital: condyle: absent, notochordal = 0 present = 1 **\$165; #5**
- Basipterygoid junction: basipterygoid process fits into socket recessed into epipterygoid = 0, pterygoid/epipterygoid forms narrow bar and clasps basipterygoid process fore and aft = 1 #6
- 7. Exoccipitals: meet skull table: absent = 0, present =1 #7
- 8. Exoccipital contributes to condyle: absent = 0, present = 1 #8
- 9. Exoccipitals enlarged to form double horizontally orientated occipital condyle, (may exclude basioccipital from articular surface): absent = 0, present = 1 similar to \$161;
 #9
- 10. Frontal parietal length ratio: frontals shorter = 0; longer = 1, subequal = 2 similar to \$35; #10
- 11. Frontal anterior margin wedged between nasals: absent = 0, present = 1 similar to\$38; #11
- 12. Frontal nasal length ratio: frontals approximately equal to or less than one-third as long as nasals = 0, more than one-third as long = 1 similar to \$13; #12
- 13. Intertemporal present: present = 0, absent = 1 **\$60; #13**

- 14. Intertemporal smaller than supratemporal = 0, or larger than/comparable in size with supratemporal = 1 modified from %45
- 15. Intertemporal lateral edge: not interdigitating with cheek = 0, interdigitates = 1 \$61;#14
- 16. Intertemporal contacts squamosal: absent = 0, present = 1 **\$62; #15**
- 17. Jugal deep below orbit (vs narrow process): 50% -> 50% orbit diam = 0, <50% = 1
 \$91; #16
- 18. Jugal contribution to orbit margin: less than one-third = 0, equal to or more than one-third = 1 Δ 7
- 19. Jugal alary process on palate: absent = 0, present = 1 \$90; #17
- **20.** Jugal length of postorbital region relative to one-third of the length of the postorbital cheek region: greater = 0 or less =1 **not previously used**
- 21. Jugal extends anterior to anterior orbit margin: absent = 0, present = 1 \$94; #18
- 22. Jugal not interposed between maxilla and quadratojugal thus not contributing to skull lower margin = 0 or interposed = 1 = #21; @27
- 23. Jugal V-shaped indentation of posterodorsal margin: absent = 0, present = 1 \$93
- 24. Lacrimal contributes to narial margin: absent, excluded by anterior tectal = 0: present = 1, absent, excluded by nasal/maxillary or prefrontal/maxillary suture = 2 #19; %244
- 25. Lacrimal reaches orbit margin (= prefrontal/ jugal suture): present = 0, absent = 1
 =\$24; #20
- 26. Maxilla sutures to vomer: absent = 0, present = 1 **\$121; #22**
- 27. Maxilla external contact with premaxilla: narrow contact point not interdigitated = 0, interdigitating suture = 1 #23; %245
- 28. Maxilla highest point in posterior half = 0, anterior third of its length = 1, or at its midlength = 2 @44
- 29. Maxilla extends behind level of posterior margin of orbit: present = 0, absent = 1 modified from \$98; #24
- 30. Maxilla sutures to prefrontal: absent = 0, present = 1 \$21
- 31. Maxilla premaxilla contact shelf-like mesial to tooth row on palate: absent = 0, present = 1 \$7; #31

- 32. Median rostral (=internasal): mosaic = 0, paired = 1, single = 2, absent = 3 #25
- 33. Nasals contribute to narial margin: absent = 0, present = 1 #26; 247
- 34. Nasal parietal length ratio less than 1.45 = 0 or greater than 1.45 = 1 = **\$15; #27**
- 35. Nasal smaller in area than postparietal: absent = 0, present = 1 #28
- 36. Opisthotic paroccipital process ossified and contacts tabular below post-temporal fossa: absent = 0, present = 1, post-temporal fenestra absent = 2 #29; in part %81
- 37. Opisthotic forms substantial plate (with supraoccipital if present) beneath skull table, separating it from the exoccipitals: present = 0, absent = 1 \$168; #30
- 38. Parietal meets tabular: absent = 0, present = 1 **\$39; #31**
- 39. Parietal postorbital suture: absent = 0, present = 1 \$40; #32
- 40. Parietal anterior portion extent relative to orbit midlength: in front of = 0, level with = 1, posterior to = 2 **\$41; #33**
- 41. Parietal shape of anteriormost third: not wider than frontals = 0, at least marginally wider = 1 \$42; #34
- 42. Parietal postparietal suture strongly interdigitated: absent = 0, present = 1 \$45;
 #36
- 43. Postfrontal prefrontal contact: broad = 0; or point-like = 1 @10
- 44. Postfrontal prefrontal suture: anterior half of orbit = 0, middle or posterior half of orbit = 1, absent = 2 #43; @9
- 45. Postorbital suture to skull table (intertemporal or supratemporal) interdigitating vs smooth: smooth = 0, interdigitating = 1 #37
- 46. Postorbital without distinct dorsomedial ramus for postfrontal = 0, with incipient ramus = 1, with elongate ramus = 2. @14
- 47. Postorbital shape: irregularly polygonal = 0, broadly cresentic and narrowing to a posterior point = 1 \$78; #38
- 48. Postorbital longer than anteroposterior width of orbit: absent = 0, present = 1 not previously used
- 49. Postorbital at least one quarter of the width of the skull table at the same transverse level: absent = 0, present = 1 \$81; #40
- 50. Postparietal: longer than wide = 0, approximately square or pentagonal = 1, wider than long = 2 similar to \$49 but split; #41

- 51. Postparietal occipital flange exposure: absent = 0, present = 1 reworded from \$52;#42
- 52. Postparietal exoccipital suture: absent = 0, present = 1 \$51
- 53. Prefrontal less than three times longer than wide: present = 0, more than, = 1 \$16; %13
- 54. Prefrontal enters naris: absent = 0, present = 1 **\$20**
- 55. Prefrontal contributes to half or more than half anteromesial orbit margin = 0, less than half = 1 22
- 56. Premaxilla posterodorsal alary process onto snout: absent = 0, present = 1 \$1; #44
- 57. Premaxilla forms part of choanal margin: broadly = 0, point = 1, not, excluded by vomer = 2 #45; %251
- 58. Preopercular present = 0, absent = 1 **\$99; #46**
- 59. Squamosal posterodorsal margin shape: convex = 0, sigmoid or approximately straight = 1, entirely concave = 2 similar to \$84 but split; #47
- 60. Squamosal contact with tabular: smooth = 0, interdigitating = 1, absent = 2 similar to \$71 but split; #48
- 61. Squamosal suture with supratemporal position: within skull table = 0, at apex of temporal embayment = 1, dorsal to apex = 2, ventral to apex = 3 #49; %252
- 62. Squamosal anterior part lying behind mid-parietal length: present = 0, absent = 1\$83; #50
- 63. Squamosal interdigitating suture with supratemporal: absent = 0, present = 1 \$66;#53
- 64. Squamosal contacts tabular on dorsal surface: absent = 0, present = 1 \$70, %53
- 65. Supratemporal present as a separate ossification: present = 0, absent = 1 \$63; #51
- **66.** Supratemporal forms part of skull margin posteriorly: absent = 0, present = 1 #52
- 67. Tabular lateral horn (subdermal unornamented component): absent = 0, button = 1, blade = 2 similar to \$68+69; #54
- 68. Tabular prolonged posterolateral ornamented surface absent = 0, present = 1; @17
- 69. Tabular emarginated lateral margin: absent = 0, present = 1 #55; %253

70. Tabular occipital flange exposure: absent = 0, extends as far ventrally as does postparietal = 1, extends further ventrally than does postparietal = 2 similar to £74 but split; #57

Palate

- 71. Ectopterygoid as long or longer than palatines: present = 0, absent = 1 \$137; #58
- 72. Ectopterygoid reaches subtemporal fossa: absent = 0, present = 1 #59; %256 (corrected: subtemporal fossa, not adductor fossa)
- 73. Ectopterygoid palatine exposure: more or less confined to tooth row = 0, broad mesial exposure additional to tooth row = 1 #60; %257
- 74. Lateral rostral present: present = 0, absent = 1 **\$9 %6**
- 75. Parasphenoid grooved ventrally about half of length = 0, vs narrow V-shaped section cultriform process along whole length = 1, flat and more or less broad = 2 #65
- 76. Parasphenoid cultriform process shape: biconvex = 0, narrowly triangular =1, parallel-sided = 2, or with proximal constriction followed by swelling = 3 modified from Δ 47
- 77. Parasphenoid depression in body: absent = 0, single median = 1, double = 2

\$171+172; #66

- 78. Parasphenoid posterolateral wings (ridged): absent = 0, present = 1 \$170; #67
- 79. Parasphenoid wings: separate = 0, joined by web of bone = 1 #68; 260
- 80. Parasphenoid contacts or sutures to vomers: present = 0, absent = 1 #69; 261
- 81. Parasphenoid carotid grooves: curve round basipterygoid process = 0, lie posteromedial to basipterygoid process (or enter via foramina there) = 1, absent = 2 #70; 262
- 82. Parasphenoid/basisphenoid ventral cranial fissure: not sutured = 0, sutured but traceable = 1, eliminated = 2 \$9; #71
- 83. Pterygoids separate in midline = 0, meet in midline anterior to cultriform process = 1 \$145; #61
- 84. Pterygoids flank parasphenoid for most of length of cultriform process = 0, not so = 1 #62

- 85. Pterygoid quadrate ramus margin in adductor fossa: concave = 0, with some convex component = 1 =\$143; #63
- 86. Pterygoids not visible in lateral aspect below ventral margin of jugal and quadratojugal = 0, or visible = 1 #64; @50
- 87. Pterygoid junction with squamosal along cheek margin: unsutured = 0, half and half= 1, sutured entirely = 2 #64
- 88. Vomers separated by parasphenoid > half length: present = 0, absent = 1 #72
- 89. Vomers separated by pterygoids: for > half length = 0, < half length = 1, not separated = 2 #73
- 90. Vomer contributes to interpterygoid vacuity: absent = 0, present = 1 \$120; #87
- 91. Vomers as broad as long or broader = 0, about twice as long as broad or longer = 1modified from \$117; #74

Upper Dentition

- 92. Ectopterygoid fang pairs: present = 0, absent = 1 #76; \Diamond 30
- 93. Ectopterygoid row (3+) of smaller teeth: present = 0, absent = 1 \$138; #77
- 94. Ectopterygoid denticle row lateral to tooth row: present = 0, absent = 1 modifiedfrom #78
- 95. Ectopterygoid / palatine shagreen field: absent = 0, present = 1 \$136; #79
- 96. Maxilla tooth number: > 40 = 0, 30-40 = 1, < 30 = 2 **#80;** \Diamond **33**
- 97. Maxillary caniniform teeth (about twice the size of neighbouring teeth): absent = 0, present = 1 **#81**
- 98. Palatine fang pairs: present = 0, absent = 1 \$127; #82
- 99. Palatine row of smaller teeth: present = 0, absent = 1 \$130; #83
- 100. Palatine denticle row lateral to tooth row: present = 0, absent = 1 modified from #84
- 101. Parasphenoid shagreen field: present = 0, absent = 1 #85; %270
- 102. Parasphenoid shagreen field anterior and posterior to basal articulation = 0, posterior to basal articulation only = 1, anterior to basal articulation only = 2 #86;
 %271

- 103. Pterygoid shagreen: dense = 0, a few discontinuous patches or absent = 1 #87;
 %272
- 104. Premaxillary teeth with conspicuous peak: absent = 0, present = 1 #89; %274
- 105. Premaxillary tooth number: > 15 = 0, 10 14 = 1, < 10 = 2#90
- 106. Vomer fang pairs: present = 0, absent = 1 **\$118; #91**
- 107. Vomerine fang pairs noticeably smaller than other palatal fang pairs: absent = 0, present = 1 #92; $\diamond 38$
- 108. Vomer anterior wall forming posterior margin of palatal fossa bears tooth row meeting in midline: present = 0, absent = 1 =\$122; #93
- 109. Vomerine row of small teeth : present = 0, absent = 1 # 74; modified from @60
- 110. Vomerine shagreen field: absent = 0, present = 1 **\$119; #95**
- 111. Vomerine denticle row lateral to tooth row: present = 0, absent = 1 #96; %279
- 112. Vomer with toothed anterolateral crest: present = 0, absent = 1 \$122; %89
- 113. Upper marginal teeth number: greater than lower = 0, same = 1, smaller than lower = 2 =\$221 but split; #97

Lower jaw characters

- 114. Adductor fossa faces dorsally = 0, mesially = 1 \$217; #98
- 115. Angular mesial lamina interdigitating suture with prearticular: absent = 0, present
 = 1 \$195; #99
- 116. Angular reaches posteriormost point of lower jaw: absent = 0, present = 1 \$197;
 #100
- 117. Coronoid (anterior) contacts splenial: absent = 0, present = 1 \$189; #101
- 118. Coronoid (anterior) contacts postsplenial: absent = 0, present = 1 #102
- 119. Coronoid (middle) contacts postsplenial: absent = 0, present = 1 \diamond 47
- 120. Coronoid (middle) separated from splenial: present, by prearticular = 0, absent = 1, present, by postsplenial = 2 @98 (in part); 046
- 121. Coronoid (posterior) posterodorsal process: absent = 0, present = 1 **\$214; #103**
- 122. Coronoid (posterior) posterodorsal process visible in lateral view: absent = 0, present = 1 \$215; #104

- 123. Coronoid: at least one has fang pair recognisable because at least twice the height of coronoid teeth: present = 0, absent = 1 \$203+204+211/+213; #105
- 124. Coronoid: at least one has fangs recognisable because noticeably mesial to vertical lamina of bone and to all other teeth: present = 0, absent = $1 \diamond 11$
- 125. Coronoid: at least one has organised tooth row: present = 0, absent =1 modified from \$205
- 126. Coronoid: at least one carries shagreen: absent = 0, present = 1 modfied from \$204; #106
- 127. Coronoid with a row of very small teeth or denticles lateral to tooth row: present= 0, absent = 1 not previously used
- 128. Coronoid: size of teeth (excluding fangs) on anterior and middle coronoids relative to dentary tooth size: about the same = 0, half height or less = $1 \diamond 77$
- 129. Dentary with parasymphysial fangs internal to marginal tooth row: present = 0, absent = 1 #107
- 130. Dentary tooth number: more than 70 = 0, 56-70 = 1, 46-55 = 2, 36-45 = 3, less than 35 = 4 **not previously used**
- 131. Dentary with a row of very small teeth or denticles lateral to tooth row: present = 0, absent = $1 \diamond 81$
- 132. Dentary external to angular + surangular, with chamfered ventral edge and absent interdigitations: absent = 0, present = 1 \$184; #107
- 133. Dentary ventral edge: smooth continuous line = 0, abruptly tapering or 'stepped' margin = 1 *17
- 134. Mandibular sensory canal: present = 0, absent = 1 *46; #109
- 135. Mandibular canal exposure: entirely enclosed apart from pores = 0, mostly enclosed = 1, mostly or entirely open = 2 \$116; #110
- 136. Mandibular oral sulcus/ surangular pit line: present = 0, absent = 1 #111; =048
- 137. Meckelian bone visible between prearticular and infradentary series: present = 0, absent = 1 #112; %282
- 138. Meckelian bone or space exposure in middle part of jaw, depth much less than prearticular = 0, depth similar to prearticular = $1 \diamond 26$

- 139. Meckelian foramina/ fenestrae, dorsal margins formed by; Meckelian bone = 0, prearticular = 1, infradentary (postsplenial) = $2 \diamond 25$
- 140. Adsymphysial tooth plate: present = 0, absent = 1 %178; #113
- 141. Adsymphysial plate fang-pair (distinct from other teeth): absent = 0, present = 1\$179; #114
- 142. Adsymphysial plate dentition: shagreen, denticles or irregular tooth field = 0, organised dentition aligned parallel to jaw margin = 1, no dentition = 2 \$180+181; modified from ◊30
- 143. Adsymphysial lateral foramen present: absent = 0, present = $1 \diamond 28$
- 144. Adsymphysial mesial foramen present: absent = 0, present = 1 \Diamond **29**
- 145. Postsplenial with mesial lamina: absent = 0, present = 1 **\$192; #115**
- 146. Postsplenial pit line present: present = 0, absent = 1 133; 48
- 147. Postsplenial suture with prearticular: absent = 0, present but interrupted by Meckelian foramina or fenestrae = 1, uninterrupted suture = 2 \diamond 35
- 148. Prearticular shagreen field, distribution: gradually decreasing from dorsal to ventral = 0, well defined dorsal longitudinal band = 1, scattered patches or absent = $2 \diamond 42$
- 149. Prearticular sutures with surangular: absent = 0, present = 1 \Diamond **39**
- 150. Prearticular with longitudinal ridge below coronoids: absent = 0, present = 1 \Diamond 37
- 151. Prearticular centre of radiation of striations: level with posterior end of posterior coronoid = 0, level with middle of adductor fossa = 1, level with posterior end of adductor fossa = $2 \diamond 36$
- 152. Splenial, rearmost extension of mesial lamina closer to anterior margin of adductor fossa than to the anterior end of the jaw: absent = 0, present = 1 \$188; #116
- 153. Surangular crest: absent = 0, present = 1 #117; \diamond 44

General skull characters

- 154. Skull longer than broad = 0, as broad as long = 1, or broader than long = 2 @3
- 155. Preorbital region of skull less than twice as wide as long = 0, or at least twice as wide as long = 1 @4
- 156. Anterior palatal fenestra: single = 0, double = 1, absent = 2 \$159; #118

- 157. Internarial/ interpremaxillary fenestra (independent of presence of median rostrals): absent = 0, present = 1 \$102; #119
- 158. Interorbital distance compared with maximum orbit diameter: greater = 0, smaller
 = 1, subequal = 2 \$120
- 159. Interpterygoid vacuities: absent = 0, at least 2 x longer than wide = 1, < 2 x longer than wide = 2 \$154+156; #122
- 160. Naris position: ventral rim closer to jaw margin than height of naris = 0, distance to jaw margin similar to or greater than height of naris = 1 #123; %283
- 161. Naris shape: slit-like = 0, round or oval = 1, upper margin ragged = 2 #124;
 %284
- 162. Naris shape: ventrally facing = 0, dorsolaterally facing = 1 #125; %285
- 163. Orbit shape: round or oval = 0, angle at anteroventral corner = 1, angle at posteroventral corner = 2: emarginated margin including jugal, lacrimal and prefrontal = 3 modified from \$105; #127
- 164. Orbit position re snout/postparietal length: centre closer to front than rear = 0, centre near middle = 1, centre closer to rear than front = 2 # 128; **(a)**1
- 165. Orbit position re snout /quadrate length: centre closer to front than rear = 0, centre near middle = 1, centre closer to rear than front = 2 #129
- 166. Pineal foramen position along interparietal suture: behind midpoint = 0, at the midpoint = 1, anterior to midpoint = 2 \$107; #130
- 167. Suspensorium proportions: quadrate to anterior margin of temporal embayment about equal to maximum orbit width (discounting any anterior extensions) = 0, quadrate to anterior margin of temporal embayment < maximum orbit width = 1, quadrate to anterior margin of temporal embayment > maximum orbit width = 2 #132; %287
- 168. Skull table/cheek junction: smooth profile = 0, square/ abrupt profile = 1 # 133
- 169. Skull table shape: longer than broad = 0, approximately square = 1, shorter than broad = 2 #134; @3
- 170. Ornament character: regular, dense, but no star-burst pattern = 0, fairly regular pit and ridge with star-burst pattern at regions of growth = 1, irregular but deep = 2, irregular but shallow = 3, absent or almost absent = 4 #135; %288

Postcranial characters

- 171. Centra: intercentrum dominant = 0, pleurocentrum dominant = 1, holospondylous
 = 2 modified from #136 & %289
- 172. Centra strongly notochordal such that notochordal space more than 2/3 diameter of entire centrum: present = 0, absent = 1 **not previously used**
- 173. Centra (trunk) pleurocentra fused midventrally: absent = 0, present = 1 \$293;
 #137
- 174. Centra (trunk) pleurocentra fused middorsally: absent = 0, present = 1 \$295;
 #138
- 175. Centrum (sacral) not distinguishable by size or shape from pre- and postsacrals = 0, distinguishable = 1 #139; %290
- 176. Clavicles meet anteriorly: present = 0, absent = 1 \$228; #140
- 177. Cleithrum co-ossified with scapulocoracoid = 0, separate = 1 \$227; #42
- 178. Cleithrum smoothly broadening to spatulate dorsal end = 0, distal expansion marked from narrow stem by notch or process or decrease in thickness = 1, end simply tapering = 2 #142; %291
- 179. Cleithrum stem cross section at mid section, flattened oval = 0, complex = 1, single concave face = 2 #143; %292
- 180. Humerus ends more or less untorted = 0, ends offset by > 60 degrees = 1 #152;
 %294
- 181. Humerus L-shaped = 0, waisted but no shaft = 1, with distinct and slender shaft = 2 \$247 split; #143
- 182. Humerus accessory foramina present = 0, absent = 1 \$254; #154
- 183. Humerus latissimus dorsi process part of ridge = 0, distinct but low process = 1, spike = 2 #155; %295
- 184. Humerus latissimus dorsi process position compared with deltopectoral crest:more proximal to head = 0, equidistant from head = 1 #156
- 185. Humerus latissimus dorsi process position relative to ectepicondyle: offset anteriorly = 0, in line = 1 \$238; #157

- 186. Humerus latissimus dorsi process confluent with deltopectoral crest: present = 0, distinct from = 1 \$241; #158
- 187. Humerus anterior margin: smooth finished bone convex margin = 0, anterior keel with finished margin = 1, cartilage-finished = 2, smooth concave margin = 3 **#159**
- 188. Humerus radial facet position: distal and terminal = 0, anteroventral = 1, ventral
 = 2 \$248 split; #160
- 189. Humerus radial/ulnar facets: confluent = 0, separated by perichondral strip of bone = 1 similar to \$239; #162
- 190. Humerus with distinct supinator process: absent = 0, present = 1 \$240; #163
- 191. Humerus with ventral humeral ridge: present = 0, absent = 1 \$244; #164
- 192. Humerus ectepicondyle distinct: present = 0, absent = 1 **\$246**; **#165**
- 193. Humerus ectepicondylar ridge distal end aligned with ulnar condyle = 0, between radial and ulnar condyles = 1, aligned with radial condyle = 2 \$246; #165
- 194. Humerus entepicondyle width relative to half humeral length: greater = 0, less = 1 \$252; #166
- 195. Humerus entepicondyle width relative to humeral head width: smaller = 0, greater = 1 \$253; #167
- 196. Interclavicle body shape (distinguished from parasternal process): rhomboid, longer than broad = 0, broader than long = 1 \$231+232; #169
- 197. Interclavicle parasternal process shape: absent or tapering = 0, parallel sided = 1
 \$230; #170
- 198. Neural arch ossification: paired in adult = 0, single in adult = 1 #171; %298
- 199. Neural arch (atlas) halves fused: absent = 0, present = 1 #172
- 200. Neural arches with distinct convex lateral surfaces ('swollen'): absent = 0, present = 1 #174; %220
- 201. Neural arches of trunk vertebrae fused to centra: absent = 0, present = 1 \$296;#175
- 202. Radius: longer than ulna = 0, same length as ulna = 1, shorter than ulna (including olecranon process if present) = 2 #178; %186
- 203. Ribs (trunk): straight = 0, ventrally curved = 1 **\$280; #179**

- 204. Ribs (trunk) not longer than height of neural arch plus centrum = 0, less than 2.5 x height of neural arch plus centrum = 1, more than 2.5 x height of neural arch plus centrum = 2 #180; %302
- 205. Ribs (trunk) tapered distally or parallel-sided = 0, expanded distally into overlapping posterior flanges = 1 modified from \$282; #181
- 206. Ribs (trunk) bear proximodorsal (uncinate) processes: absent = 0, present = 1 modified from \$281; #182
- 207. Ribs (trunk) differ strongly in length and morphology along 'thoracic' region:absent = 0, present = 1 #183; %305
- 208. Ribs (cervical): flared distally = 0, tapered distally = 1 #184
- 209. Scapulocoracoid dorsal blade: absent = 0, present = 1 #187; %308
- 210. Scapular ossification separate from coracoid: absent = 0, present = 1 \$233; %188
- 211. Gastralia: tapered and elongate, 4 or >4 x longer than broad = 0, ovoid = 1, around 3 x longer than broad one end tapering = 2 #189; %309
- 212. Pelvis: illium, ischium, pubis not separate ossifications = 0, separate = 1 # 176
- 213. Illium: post illiac process and dorsal blade present = 0, only post iliac process present = 1 modified from \$259; #165

#NEXUS begin data; dimensions ntax=45 nchar=213; format missing=? symbols="0~4"; matrix Acanthostega Asaphestera 1?1?101112001???11?0010101111003100?110201111000001100002111?0?11000011?11?????? 110221?1232111??1??121????3?????1?101?01?12000?10?10 Balanerpeton 2201102121001111001101101001111101111112??11111?141001?11?210?001122?0?00?12012? 11000210210?00?112??11????3???1??1?00100010100001?111 Baphetes 01???00??1100010?100?100?001000111010002010?110?010000?01102311001100001111120?1 Diploradus Ossirarus Caerorhachis ?????0??????00?000???00??1?1000?1????0021101111?02??0??011??111?0?0?10001??210?? Casineria Colosteus 1????0???1111???11?00?02010101030110?0121100100111??1110110000110000001?1112???1 ?21100?1100001010011??0011?100110110????0???0?1?0411002?111?????0???1??00010010 Crassigyrinus 1??1?00??20000000?0000101110013100??00211110000010001002112110001111?000111110122?00011000001010001020000001010101?0000??11111?141010211010010001021011000111012111002102?01??11?10001111200010?11000000112000???010 Dendrerpeton 1????1???00000101100100101110003100??0021100101002??0001?122201001001???111200?0 11011101210000?111?111????32?010?1?00??00211000010111 Doragnathus Discosauriscus 0111?001020101111101010100011003101??10211111010021000000122211001100201112?0111

110012112311111012?0011001300010?1?111110211001011110Edops 011111110100001000010001011000300011002010010111211000?11?2?1100?101101111200?1 02111?212000111200110?0020111111211?1???1?11111?0310?1?1??????0011??10???0020011 Eoherpeton ????????100000011?010011?0100?3100??10211?10000021?0000?1121?0000100?0?01??111? 11011211221110??1??0010011010110201??1?00?1200??1??10 Eryops 0?11111111001???00011000101100031101100101001000021100111122111000010101112300?? 020110212100111110111?011011111111111021111111?021011?11?1???00112210?0100200111102202121010011112111101320110201101000201011010?11 Eucritta 1???????200001010?000010?010003?00??0021110100002??000??112201001000????1?110?1 Eusthenopteron 0?00012000000001?00?0????000000010??0?0000000100?0? Gephyrostegus 221011?12010111100111?00201111111111110111101111?031001?11011????1112?0?110020200 1102121114111011122?01????0????0???1?100021200101?010 Greererpeton Hyloplesion 1???????0011???1000010100011003110??1121110100002??00101101?10110000?11?12100?0 110002?12421111?12??2????321?11???101100212000110?11 Ichthyostega 001000002001???000000001010012000??0000100100001?000002001001000001?111??100?1 00011010110000?1000000?????11?00001110?00212111?00000 Loxomma 01???0???2000010??00?000?00?000?10011002110?111?010000?0110231100110000?111120?? 02???????0011020011??0020011011?11?????1?????1???????021101?????111?1?1?000?0??0 Metaxygnathus Megalocephalus 0?11100101001?????10?002?001000111010002010?100?020000?01102311001100001111120?1 02100021200011110011000020011111011010000?11101?1410102110201100112210?000000?00Microbrachis ??11101110011???1000010?00?110031100111211011001021100100101?0?11?000111?11100?1 12011000001111120111000021?110111111100010111111?141010211011????112200?110120010110002?1212111?112??11????3???11???10011021200011?111 Occidens

Ossinodus ?????0????????????????010000?0?0100????0?201211011?00120???0??012111001000?0111?????? Paleothyris 1?11100001011???10?1010100011003100?011211121000021000100112?0?0010002???1?120?1 11011211131111101??121???32011021?1110012120001102?? Panderichthys Pederpes 1????0??????01?111??0??10?00000?1????00??100?100021?0010?0123010011002010?03?101 21221?212100000111100121012??110?0?000000102111011210 Pholiderpeton $021010011010010010111?\\ 0021?11011111100?1111111?\\ ?310102011100200111211?\\ 110020210$ 2102210112111111111?0????????1011?0010002120000102?? Proterogyrinus 1?111001010001001100?011??0100300000010211100000021000?0?11211000121120001131111021111?????10100001101002??????11101000??0??1??021011?1111??????1?1100?0210 110212211311101111?001111122?110001001000212000010210 Seymouria 111112012111111011111111?1321111?10111111211101011?10 Sigournea Silvanerpeton 1?1110???200010012?001010??11003100??102?111001001??0010?110110100200???11?110?1 11011211031010?111?001????1????0?1000?000212000010210 Tiktaalik 20000??0?0????????????????????????????00010000?0100?0??100????0000????0?00????000?0200 ?012102020?????10?000????001000110??????00?100??0?0? Tulerpeton ????????100???0111101101100111010100????21????10010 Ventastega Westlothiana ?????0???1011???10??0???0?01?0?3100??112111010100??????0?102?10000000????1??10?1 11011111241111??1??121???32?11121???1?01212000110?11 Whatcheeria ?????00??110001111?110011?0100?3110100010100100001100000?0121110011012??????10?
Koilops Ymeria Perittodus Aytonerpeton ;

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end;
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begin trees;
tree tnt_1 = [&R]
(18,(30,(24,(37,(1,((28,(6,(((5,(10,(12,((23,((17,((19,(36,((29,(40,(2,(21,26)))
),(8,(13,34))))),(15,(32,33)))),(7,((14,(11,(3,16))),((42,(38,(9,20))),(35,45)))
))),(4,25))))),(27,41)),(31,(43,44))))),(22,39))))));
```

end;

Supplementary Table 1. Table 1a shows the localities, stratigraphy, age, spore zone, environment and number of samples from each site. Table 1b shows the fusinite abundances as a percentage of total phytoclasts taken for each stage sampled.

Locality/references	Stratigraphy	Age	Spore	Environment	Number of
			Biozone(s)		samples
1 ^{76,77}	Pathead Fm	Late Viséan	VF	Deltaic –	8
				nearshore marine	
1 ^{76,77}	Sandy Craig Fm	Mid-Viséan	NM – VF	Fluvio-deltaic	3
1 ^{76,77}	Pittenweem Fm	Mid-Viséan	NM	Deltaic –	14
				nearshore marine	
1 ^{76,77}	Anstruther Fm	Early – mid-	тс	Deltaic –	13
		Viséan		nearshore	
				marine ²	
2 ^{77,78}	Fife Ness Fm	Earliest Viséan	Pu – TS	Fluvio-deltaic	4
3 ^{15,78}	Ballagan Fm	Tournaisian	CM	Fluvio-lacustrine	12
4 ^{15,78}	Ballagan Fm	Tournaisian	VI – CM	Fluvio-lacustrine	61
5 ^{79,80}	Stensiö Bjerg	Latest	$LL - LN^5$	Fluvio-lacustrine	9
-	Fm	Famennian			

Extended data table 1a

Localities: 1, Anstruther to St Monans coastal sections, East Fife, Scotland; 2, Fife Ness, Scotland; 3, 'Willie's Hole', Chirnside; 4, Burnmouth Shore; 5, Celsius Bjerg, East Greenland.

Extended data table 1b

Famennian	Tournaisian		Viséan
2.4	Burnmouth	Willie's Hole	4.8
1.8	5.2	2.8	0.8
1.2	4.0	2.4	1.2
0.2	0.6	3.0	6.0
4.2	4.8	3.0	1.8
2.4	1.4	0.6	1.2
1.6	1.6	1.6	2.8
0.8	5.2	2.4	1.0
5.0	4.6	0.6	1.4
	3.0	2.3	2.0
	2.8	3.6	0.4
	0.2	1.2	2.4
	3.4	0.2	2.4
	0.2		1.6
	0.2		2.2
	1.0		2.4
	3.4		4.2
	1.2		0.4
	2.2		2.6
	1.0		2.6

2.8	3.8
0.6	9.0
2.0	2.6
2.0	0.6
2.0	1.6
2.2	1.4
4.2	2.6
5.2	0.2
1.2	0.8
2.2	3.0
1.2	4.2
1.8	2.4
2.2	2.6
1.8	0.4
0.4	3.8
0.8	 3.4
0.4	3.2
4.2	5.4
5.8	4.2
4.6	2.2
1.0	4.8
0.4	4.2
0.4	
1.0	
1.2	
0.6	
1.8	
9.4	
3.8	
0.4	
2.0	
2.0	
1.6	
3.4	
2.2	
3.4	
5.6	
1.8	
0.6	
0.6	
2.2	
1.4	

Famennian – Viséan Fusinite abundance (% total phytoclasts)