

25 uniform pumping rate and (b) under a transient regime. The analysis is performed within a
26 numerical Monte Carlo framework. Our results show that contaminant concentration
27 breakthrough curves (BTCs) at the well are markedly affected by the transient pumping
28 strategy according to which the well is operated. Our results document the occurrence in time
29 of multiple peaks in the mean and variance of flux-averaged concentrations at the extraction
30 well operating at a transient rate. Our findings suggest that lowest and largest values of mean
31 and variance of flux-averaged concentration at the well tend to occur at the same time. We
32 show that uncertainty associated with detected BTCs at the well increases for pumping regimes
33 displaying a high degree of temporal variability. As such, the choice of the type of engineering
34 control to the temporal sequence of pumping rates could represent a key factor to drive
35 quantification of uncertainty of the contaminant concentration detected at the well. It is
36 documented that pumping rate fluctuations induce a temporally oscillating risk pattern at the
37 well, thus suggesting that the selection of a dynamic pumping regime has a clear influence on
38 the temporal evolution of risk at the well.

39

40

41 **1. Introduction**

42 Pumping wells are widely employed for groundwater supply in the context of urban,
43 agricultural, and industrial activities. Water management agencies typically schedule
44 groundwater extraction through a well-defined temporal sequence of pumping rates to
45 maximize benefits to anthropogenic activities and minimize the environmental footprint of the
46 withdrawal process. The temporal dynamics of well pumping operations induce changes in the
47 subsurface flow field which in turn can affect spreading and mixing of contaminant plumes. A
48 key challenge is to provide estimates of contaminant concentration at an environmentally
49 sensitive location while accounting for the joint effect of the temporal dynamics of an operating
50 pumping well and the (typically uncertain) spatial variability of the hydrogeological properties
51 of the subsurface reservoir within which a potentially harmful solute is residing and migrating.

52 The effect of the spatial variability of hydraulic properties of aquifers, such as hydraulic
53 conductivity, on contaminant transport has been broadly studied (e.g., *Gelhar, 1993; Dagan*
54 *and Neuman, 1997; Rubin, 2003*). Due to the high cost of data acquisition campaigns and
55 limited financial resources, a highly detailed characterization of the spatial distribution of
56 hydrogeological attributes (e.g., hydraulic conductivity) of an aquifer system at the field scale
57 of interest is virtually unfeasible. As a consequence, this lack of knowledge leads to uncertainty
58 associated with predicted values of contaminant concentration at environmentally sensitive and
59 strategically valuable locations, such as pumping wells. As noted in the following, while a
60 considerable amount of published works address the analysis of contaminant transport under
61 steady-state pumping, the way a given time-varying pumping schedule affects contaminant
62 plume behavior in heterogeneous aquifers is tackled only marginally.

63 Uncertainty in the quantification of the extent of pumping well protection regions (both
64 in terms of well catchment and time-related capture zones) in the presence of random spatial
65 distributions of aquifer parameters and under the assumption of steady-state background
66 groundwater flow to which one or more wells extracting a constant flow rate are superimposed
67 has been addressed, amongst others, by *Cole and Silliman (1997)*, *Varljen and Shafer (1991)*,
68 *Franzetti and Guadagnini (1996)*, *Vassolo et al. (1998)*, *van Leeuwen et al. (1998, 2000)*,
69 *Guadagnini and Franzetti (1999)*, *Riva et al. (1999)*, *Feyen et al. (2001)* through numerical
70 Monte Carlo simulations (unconditional or conditional on data). *Stauffer et al. (2002)*
71 performed a theoretical analysis based on first-order approximation and a Lagrangian approach
72 to estimate the uncertainty associated with the delineation of well catchments in the presence
73 of a spatially variable random hydraulic conductivity field. *Lu and Zhang (2003)* used a
74 Lagrangian framework for the study of time-related capture zones in spatially randomly
75 heterogeneous aquifers. *Riva et al. (2006)* used stochastic moment equations to analyze the
76 uncertainty associated with well catchments. The work of *de Barros et al. (2013)* addressed
77 the value of hydrogeological information to assess the risk associated with pollution of an
78 operating pumping well. *Molson and Frind (2012)* applied the concept of mean life expectancy
79 (i.e., the time between plume detection and plume capture by a well) and capture probability
80 to delineate well capture zones. The life expectancy approach constitutes an alternative to
81 delineate time-of-travel capture zones for wellhead protection. *Pedretti and Fiori (2013)*
82 developed an analytical solution of contaminant transport in heterogeneous porous media
83 under purely convergent flow to analyze solute breakthrough curves (BTCs) at the well.
84 *Indelman and Dagan (1999)* analyzed the spreading of a contaminant plume injected in a
85 heterogeneous formation through a well and advected by a steady-state groundwater velocity.

86 Only a limited number of works consider the impact of transient flow regimes on the
87 delineation of capture zones or wellhead protection areas (WHPA). Amongst these,
88 *Ramanarayanan et al.* (1995) used numerical models to delineate the WHPA of two municipal
89 wells under the influence of nearby irrigation wells. Their conclusion was that seasonal
90 variation in pumping rates affected the water table depth and the extent of a target time-related
91 well capture zone. In this sense, significant seasonal variations in pumping operations should
92 be explicitly embedded in a procedure for WHPA delineation. *Reilly and Pullock* (1996)
93 employed a numerical flow model and a particle tracking method to investigate the effects of
94 external stresses, i.e. recharge (that is typically cyclic over time), on the area contributing
95 recharge to the wells. They observed that the ratio between the mean travel time of the solute
96 to the well and the period of cyclic stress can provide a quantification of the importance of
97 considering the transient effects of a cyclic recharge. The authors noted that: (a) approximating
98 the system behavior through an average steady-state condition yields results which are similar
99 to those obtained by considering the actual transient pattern of recharge when the above ratio
100 is much larger than one; and (b) considering transient flow conditions is important to
101 characterize the proper behavior of solutes released in the proximity of the boundary of the
102 well. *Jacobson et al.* (2002) incorporated the effects of uncertainty in the transmissivity and
103 the magnitude and direction of the hydraulic head gradient in the steady-state analytical
104 solution obtained by *Bear and Jacobs* (1965) for capture zones delineation. These authors
105 observed that uncertainty in the magnitude of the mean regional flow propagates to the
106 uncertainty in the length (along the direction of the regional flow) of the time-dependent
107 capture zone. Otherwise, uncertainty in the mean direction of the background flow influences
108 the uncertainty in the maximum width (along a direction normal to the average regional flow)

109 of the capture zone. *Festger and Walter* (2002) analyzed the same problem through a semi-
110 analytical approach. They concluded that temporal variations in the direction of the
111 background natural hydraulic gradient typically cause an expansion of a given capture zone
112 when compared against its steady-state counterpart, the opposite being observed in some
113 specific cases (e.g., in the presence of moderate directional variations and large gradient
114 magnitudes). *Neupauer et al.* (2014) investigated the effects of aquifer heterogeneity on the
115 spreading of a plume of conservative solute which is advected within a temporally variable
116 flow field induced by engineered injection and extraction sequences. Other authors, e.g.
117 *Assouline et al.* (2006), illustrate that transient effects associated with high-frequency irrigation
118 affect solute and water regimes and residence times under partially saturated conditions.

119 Additional examples of contaminant transport studies which incorporate the effects of
120 transient pumping strategies include the works of *Chang et al.* (1992), who simulated unsteady
121 flow to optimize time-varying pumping rates for groundwater remediation; *Vesselinov* (2007),
122 who studied the influence of transient flow and of uncertainty in longitudinal and transverse
123 dispersivities on the delineation of well capture zones; *Chen et al.* (2012), who evaluated solute
124 transport in divergent radial flow fields created by multistep pumping; and *Leray et al.* (2014),
125 who detected the occurrence of non-negligible solute concentrations at a pumping well
126 following induction of a transient flow field by the instantaneous activation of the well. In
127 contrast to our work, *Leray et al.* (2014) did not investigate the uncertainty of the contaminant
128 concentrations recovered at the well.

129 Notably, in this broad context, the effects of a temporally varying pumping regime on
130 contaminant solute BTCs detected at the location of the pumping well operating in a
131 heterogeneous aquifer has not been systematically addressed. Our work is then geared towards

132 a first exploration of the joint effects of (a) random spatial variability of hydraulic conductivity
133 and (b) transient flow regime, as induced by a temporally varying pumping schedule, on this
134 key feature of contaminant transport. These results will then be employed in the context of risk
135 quantification at the well, which is considered as a sensitive strategic location with social and
136 economic value. Risk is here quantified in terms of the probability that an established threshold
137 for the concentration of a target chemical species is exceeded. As such, the objective of our
138 study is directly related to the assessment of risks associated with contamination of drinking
139 water extracted at the well (see. e.g., the works of *Hashimoto et al.*, 1982; *Frind et al.*, 2006;
140 *de Barros et al.*, 2009; and *Enzenhoefer et al.*, 2012).

141

142 **2. Problem formulation**

143 We consider fully saturated groundwater flow taking place in a heterogeneous porous
144 medium characterized by a spatially variable isotropic random transmissivity field $T(\mathbf{x})$ [L^2/T],
145 uniform porosity φ [-], constant thickness b [L] and a uniform-in-the-average natural base
146 flow Q_0 [L^3/T]. The location vector is here expressed as $\mathbf{x} = (x, y)$, in a Cartesian coordinate
147 system. Figure 1 depicts a sketch of the domain and problem we study.

148 The log transmissivity $Y(\mathbf{x}) = \ln[T(\mathbf{x})]$ is assumed to be a normally distributed and a
149 weakly stationarity correlated random field, which is fully described by its mean, m_Y , and
150 covariance function, C_Y (e.g., *Rubin*, 2003). The covariance function is statistically isotropic
151 and characterized by the variance of $Y(\mathbf{x})$, denoted by σ_Y^2 , and its (isotropic) correlation scale λ
152 [L]. A pumping well located at $\mathbf{x}_w = (x_w, y_w)$ operates with pumping rate $Q_w(t)$ [L^3/T], with t
153 [T] denoting time. Therefore, the flow field is described by the depth-averaged saturated
154 groundwater flow equation:

$$S \frac{\partial h(\mathbf{x}, t)}{\partial t} = \nabla \cdot [T(\mathbf{x})\nabla h(\mathbf{x}, t)] + Q_w(t)\delta(\mathbf{x} - \mathbf{x}_w); \quad (1)$$

155 where h [L] is the hydraulic head, S [-] corresponds to storativity and δ is the Dirac delta
 156 function. Fixed hydraulic head values are imposed on the west (left) and east (right) boundaries
 157 and no-flow boundary conditions are set on the north and south boundaries of the domain,
 158 resulting in a uniform mean base flow along the x -direction (Figure 1).

159 A tracer with initial uniform concentration c_0 [M/L³] is instantaneously injected at time
 160 t_0 [T] within a region of area A_0 [L²]. Transport of the tracer plume is described by the standard
 161 advection-dispersion equation:

$$\frac{\partial c(\mathbf{x}, t)}{\partial t} = -\mathbf{v}(\mathbf{x}, t) \cdot \nabla c(\mathbf{x}, t) + \nabla \cdot [\mathbf{D}\nabla c(\mathbf{x}, t)]; \quad (2)$$

$$c(\mathbf{x}, t_0) = c_0 \text{ for } \mathbf{x} \in A_0.$$

162 Here, c [M/L³] is solute resident concentration, $\mathbf{v} = \mathbf{q}/\varphi$ [L/T] is effective velocity (and \mathbf{q} [L/T]
 163 is the specific discharge) and \mathbf{D} [L²/T] is the local-scale dispersion tensor with longitudinal
 164 component D_x and transverse component D_y . We consider that solute transport takes place
 165 sufficiently far away from the boundaries (Figure 1). Once we obtain the space-time
 166 distribution of the concentration we evaluate the flux-averaged concentration C [M/L³] at the
 167 pumping well (*Kreft and Zuber, 1979*) and obtain the solute BTC. Details related to the
 168 computational method used to evaluate (1) and (2) are provided in Section 3.

169 Equations (1) and (2) are expressed in dimensional form. In this work, we will analyze
 170 our results in terms of relevant dimensionless groups. We start by the following functional
 171 relationship between system parameters and state variables of interest (see also, e.g., *Franzetti*
 172 *and Guadagnini, 1996*):

173

$$f_1(t, x, y, T, b, q_0, C, Q_w, \varphi, D) = 0; \quad (3)$$

174

175 with q_0 [L/T] being the imposed uniform Darcy's background groundwater specific discharge.

176 We consider q_0, Q_w, b and φ as deterministic in our analysis. The following relationship

177 between dimensionless quantities can then be written under the assumption that φ and D are

178 constant:

$$f_2\left(\frac{tq_0^2b}{Q_w}, \frac{xq_0b}{Q_w}, \frac{yq_0b}{Q_w}, \frac{T}{q_0b}, \frac{C}{C^*}\right) = 0. \quad (4)$$

179 Note that the dependence on a typical Péclet number would appear in (4) in case one would

180 consider analyzing the interplay between advective and dispersive processes in the system (see

181 Sections 3 and 4.4 for a discussion on this aspect).

182 The reference concentration C^* in (4) is taken as the contaminant concentration limit

183 value for a target chemical species. The value of the latter is established, for example, by

184 regulatory bodies/agencies. In the regulatory context, this contaminant concentration limit can

185 represent the *Maximum Contaminant Level* for drinking water. Note that Q_w in (4) indicates

186 the constant pumping well extraction rate. In this work, the constant pumping rate, which is

187 selected in the steady-state setting, coincides with the average (over the temporal window of

188 well operation) of the pumped flow rate in the transient scheme (see Section 3). Adopting the

189 constant (in time) extraction rate is then a natural choice to normalize the various terms in (4)

190 for both the steady-state and transient settings analyzed.

191 Due to the randomness of $Y(\mathbf{x})$, the feedback between a transient pumping regime and
192 contaminant breakthrough curves must be studied in a stochastic framework. Section 3
193 provides details of the computation of the flux-averaged concentration at the well.

194

195 **3. Methodology**

196 We consider a rectangular solute source, with longitudinal (along direction x) dimension
197 L_1 and transverse (along direction y) dimension L_2 , within which solute is instantaneously
198 released in the system (i.e. the area of the source is $A_0 = L_1 \times L_2$, see Figure 1). A pumping
199 well is located at a distance L (measured along the x -direction) from the contaminant source
200 and aligned with the source centroid location. The computational domain has size $\Omega = [(x, y)/$
201 $x \in (0, L_x); y \in (0, L_y)]$ with $L_x = 170$ m and $L_y = 150$ m. We considered a unit thickness aquifer
202 in all simulations.

203 In order to quantify the statistics of C , we employ a numerical Monte Carlo (MC)
204 framework. A collection of realizations of the $Y(\mathbf{x})$ field are generated using a sequential
205 Gaussian simulator (SGeMS; *Remy et al.*, 2009). The spatial covariance C_Y is modeled through
206 a stationary Exponential model, characterized by isotropic correlation scale $\lambda = 8$ m and
207 variance $\sigma_Y^2 = 3$. Consistent with previous studies (e.g., *de Barros et al.*, 2013; *Moslehi et al.*,
208 2015), a uniform generation grid, with elements of side $\Delta_x = \Delta_y = 1/8\lambda$, is employed. Transient
209 groundwater flow and contaminant transport for each realization are respectively solved
210 through the widely tested codes MODFLOW (*Harbaugh*, 2005) and MT3DMS (*Zheng and*
211 *Wang*, 1999). A Lagrangian-Eulerian method, the Method of Characteristics (MOC), is applied
212 to solve contaminant transport with MT3DMS. [Further details pertaining the computational](#)
213 [methods \(and the associated numerical errors\) adopted are documented in the literature \(*Zheng*](#)

214 *and Bennet, 2002*). A comparison between the numerical methods utilized in this work with
215 other existing codes can be found in *Hassan and Mohamed (2003)*.

216 We explore three diverse pumping scenarios, indicated as scenarios I, II and III. Each
217 scenario is characterized by a fixed total volume of water extracted (V_w) and a given operational
218 scheme according to which the pumped flow rate is varied in time. Each of the transient
219 scenarios is compared against a corresponding setting associated with steady-state pumping
220 characterized by the same total volume of water extracted from the system during the simulated
221 transient regime. Table 1 lists the main parameters used in the study. Note that the longitudinal
222 Péclet number is set to 800. The Péclet number is given by $Pe = \lambda v_0 / D_x$, where $v_0 = q_0 / \varphi$
223 denotes the imposed uniform-in-the-mean longitudinal velocity and D_x is the local scale
224 longitudinal dispersion coefficient, which we set at the constant value listed in Table 1.

225 We perform MC numerical simulations of flow and contaminant transport according to
226 the three scenarios displayed in Figure 2 (Figure 2a, b, and c respectively referring to scenarios
227 I, II and III). The selected well pumping operation schedules are inspired by patterns employed
228 in practical settings by groundwater management agencies to extract water from underground
229 resources for diverse purposes (including, e.g., drinking and irrigation). As seen in Figure 2,
230 scenario III is characterized by the highest variability of the dynamic pumping strategy.

231 Following a preliminary convergence study (not shown), we base our results on a set of
232 1000 MC simulations for each scenario and pumping strategy adopted. The results of the MC
233 simulations are post-processed with the aim of capturing the temporal evolution of the mean
234 and variance of the flux-averaged contaminant concentration recovered at the well. We note
235 that the effects of variable well pumping operations on the solute breakthrough curve at the

236 well can also depend on the ratio between the natural base flow Q_0 and the well pumping rate
237 Q_w .

238 For each scenario we quantify the risk, defined here as the probability of failure, e.g.,
239 the probability that the contaminant concentration exceeds a given critical value at the well.
240 According to this, we define risk (ξ) in this work in terms of exceeding probabilities (e.g., *de*
241 *Barros and Fiori, 2014*):

$$\xi = P[C(\mathbf{x}_w, t) \geq C_{crit}], \quad (5)$$

242 where $C_{crit} \equiv C^*$, as introduced in Section 2, and P denotes probability. Note that the risk
243 definition in Equation (5) rests on a full statistical distribution of the flux-averaged
244 concentration at the pumping well as a function of time. As shown by several authors (*Kapoor*
245 *and Kitanidis, 1998; Fiori and Dagan, 2000, de Barros and Fiori, 2014*), the mean
246 concentration is insufficient to characterize risk and higher order moments are required. The
247 mean concentration only becomes a meaningful measure of risk under ergodic conditions, i.e.,
248 large travel distances or large contaminant source dimensions with respect to the log-
249 transmissivity correlation scale. If these conditions are satisfied, macrodispersion theory is
250 expected to hold (*Dagan, 1984; Kitanidis, 1988; Chapter 10 of Rubin, 2003* and references
251 therein). In particular, *de Barros and Fiori (2014)* provides a detailed discussion on the need
252 to use of a full probabilistic distribution to quantify risk by taking into account also the
253 probability of extreme events (e.g., probability of concentration exceeding a very high or low
254 value). The exceedance probability (5) is computed directly from the numerical MC results
255 and informs decision makers on the probability that $C(\mathbf{x}_w, t) \geq C_{crit}$.

256 4. Results and Discussion

257 We start by illustrating the relevance of the problem tackled in this study through a
258 discussion of some key differences between contaminant BTCs observed at the pumping well
259 operating with a steady-state pumping strategy and a transient pumping scheme for selected
260 exemplary realizations of the $Y(\mathbf{x})$ field. Second, we present the temporal evolution of the mean
261 and variance (as well as the resulting coefficient of variation) of the flux-averaged
262 concentration at the well for the pumping strategies of scenario III. The latter is representative
263 of the main conclusions of our analysis, the results of scenarios I and II being fully documented
264 in the Supplemental Material (SM). Third, we display the temporal behavior of risk, as defined
265 in (5), for scenario III, under steady-state and transient pumping conditions. Corresponding
266 outputs for scenarios I and II are included in the SM. Finally, we demonstrate the effects of
267 local scale dispersion and aquifer heterogeneity on the uncertainty of the concentration BTCs
268 for scenario III. We show how the interplay between pumping operations, aquifer
269 heterogeneity and local scale dispersion affect the temporal patterns of the flux-averaged
270 concentration uncertainty. [The upcoming results also emphasize the importance of pumping in](#)
271 [diluting the concentration at the well.](#)

272 *4.1. Concentration breakthrough curves observed at the pumping well for different pumping* 273 *strategies*

274 In this Section, we aim to provide a qualitative illustration of the way the concentration
275 BTC at the well can be affected by the pumping operation and the heterogeneous transmissivity
276 field. A graphical depiction of the effect of a time-dependent withdrawal schedule on the

277 concentration BTC is provided in Figure 3. The latter depicts the concentration BTCs observed
278 at the pumping well for three diverse realizations of $Y(\mathbf{x})$ under the constant and temporally
279 varying pumping strategy of each scenario. Each plot within Figure 3 includes the contaminant
280 BTCs resulting from a constant and a transient extraction strategy for a given realization of the
281 $Y(\mathbf{x})$ field. Figures 3a, b, c refer to Scenario I, Figures 3d, e, f refer to Scenario II and finally,
282 Figures 3g, h, i refer to Scenario III. The three distinct realizations of the transmissivity are
283 denoted by T_1 , T_2 and T_3 (see Figures SM17-SM19 of the SM) in Figure 3 and are selected
284 solely for the purpose of illustration. These three realizations of the transmissivity fields (T_1 ,
285 T_2 and T_3) were obtained through the software SGeMS (Remy *et al.*, 2009) and are
286 characterized by the geostatistical parameters described in Section 3. As shown in Figure 3,
287 significant differences are detected between the concentration BTCs obtained with uniform
288 and variable in time pumping rates. From a qualitative standpoint, it can be noted that the
289 concentration BTCs produced by a dynamic pumping regime are characterized by peaks of
290 different strength and occurring at diverse times when compared to the corresponding BTCs
291 obtained with uniform pumping. This result has strong implications in risk analysis, inasmuch
292 as public health authorities are concerned with peak concentrations and the time of their
293 occurrence (e.g., *de Barros and Rubin, 2008; Oladyshkin et al., 2012; Siirila and Maxwell,*
294 *2012*). The results displayed in Figure 3 also illustrate the complex interaction between the
295 hydrogeological heterogeneity of the aquifer and the transient flow regime induced by a
296 temporally varying pumping schedule. These illustrative results point out the importance of
297 considering temporally varying pumping settings, as employed in practical applications, when
298 evaluating contaminant transport in a well field, with direct implication on risk caused by
299 contamination of water extracted at the well.

300 4.2. Mean and variance of solute concentration at the pumping well

301 We compute and discuss here the temporal evolution of the first two statistical moments
302 of the flux-averaged solute concentration detected at the pumping well. To streamline the
303 presentation, we only show the results related to scenario III, which encapsulates the key
304 observations of the analysis. Results associated with scenarios I and II are documented in the
305 SM.

306 Figures 4 and 5 depict the mean and the variance of the flux-averaged concentration
307 (respectively denoted as $\langle C \rangle$ and $Var[C]$) detected at the well following the adoption of a
308 uniform (Figures 4a and 5a) and dynamic (Figures 4b and 5b) pumping strategy. These results
309 indicate that the transient operational regime applied at the well has marked influence on the
310 shape of the concentration BTCs. A striking observation is that the local minimum and
311 maximum of the mean and variance of the flux-averaged concentration at the well occur at the
312 same dimensionless time for both types of pumping regimes. This observed behavior is
313 supported by theoretical developments illustrated in Appendix A.

314 Figure 4b shows that the mean flux-averaged concentration starts increasing when the plume
315 approaches the well, until a maximum value is reached between dimensionless time values of
316 0.2 and 0.3. Figure 4b also reveals that an increase in the pumping rate can lead to a sudden
317 decrease of the mean flux-averaged concentration. Due to the intermittent pumping schedule,
318 this effect is seen to extend in time causing an oscillatory pattern of the mean BTC. It can be
319 noted that activating the pumping well induces contaminant dilution at the well due to mixing
320 of contaminated and clean water. This leads to a decrease of the contaminant concentrations in

321 the water extracted at the well with respect to concentration levels [in the contaminant plume](#)
322 (*Einarson and MacKay, 2001, de Barros et al., 2009*).

323 A pronounced oscillatory pattern in time is detected for $Var[C]$ (Figure 5b). We note
324 that an increase of the pumping rate corresponds to a decrease of $Var[C]$ across the various
325 stress periods. This behavior is consistent with the observation that an increase of the extraction
326 rate enhances the likelihood that the contaminant is captured by the well. This consequently
327 tends to dampen the variability of contaminant concentrations observed at the well, with a
328 corresponding reduction of $Var[C]$. The opposite tends to happen when the pumping rate is
329 reduced (and eventually goes to zero), thus leading to increased values of $Var[C]$ at the well.

330 By comparing the variance of the flux-averaged concentration detected at the pumping
331 well operating under a temporally variable strategy for the three investigated scenarios (Figure
332 5b addresses scenario III and Figures SM4 and SM8 of the SM respectively refer to scenarios
333 I and II), we observe that the maximum values in Figure 5b (scenario III) are much higher than
334 their counterparts appearing in Figure SM4 (scenario I) and Figure SM8 (scenario II) of the
335 SM. We recall that scenario III is associated with the highest temporal variability of the well
336 operational scheme. Our results indicate that the uncertainty associated with the flux-averaged
337 concentration is higher for pumping strategy III than for pumping regimes I and II. This
338 behavior is also supported by viewing the results of Figures 6 and 7, respectively depicting the
339 coefficient of variation (CV) of the flux-averaged concentration for the constant and transient
340 pumping strategy of scenario III and comparing them against the corresponding results
341 depicted in the SM (see Figures SM9, SM10 for scenario I and Figures SM11 and SM12 for
342 scenario II). When an extraction schedule with high temporal variability (see, e.g., scenario
343 III) is applied at the well, the values of CV tend to increase, in particular during time intervals

344 when pumping is stopped. We conclude that considering a high temporal variability of the
345 pumping rate has the effect of increasing the uncertainty associated with the concentration
346 observed at the pumping well. We observe that while it is well known that aquifer
347 heterogeneity leads to uncertain contaminant transport prediction, our results indicate that the
348 engineered operational pumping schedule, manifested through $Q_w(t)$, may play a major role in
349 controlling the variability of the moments of flux-averaged concentration at the well. Results
350 in Figures 6, 7, SM9, SM10, SM11, SM12 also show that uncertainty in the solute
351 concentration at the well is large at both early and late times. This implies that the (ensemble)
352 mean is not sufficient to characterize the transport behavior and quantification of higher order
353 moments is needed.

354 *4.3. Probability of exceedance of concentration limit value*

355 We compute here the probability of failure, i.e. the probability that the contaminant
356 concentration exceeds an established limit value, C^* . This probability represents a measure of
357 risk, as defined in (5) and is computed on the basis of the MC realizations performed. The
358 exceedance probability provided in Equation (5) enables one to evaluate the probability of
359 extreme events since it contains information about the mean, variance and higher order
360 moments of the flux-averaged concentration. Figures 8 and 9 depict the temporal evolution of
361 $P(C > C^*)$ for the uniform and time-varying pumping strategies of scenario III. The choice of
362 the operational scheme (constant or variable in time pumping) to withdraw a fixed water
363 volume has a clear influence on the temporal evolution of risk (5) at the pumping well.

364 The results depicted in Figure 8 are linked to a temporal pattern which is similar to that
365 of the mean flux-averaged concentration in Figure 4a. For the cases analyzed, the results

366 suggest that higher mean concentration values are associated with increased probability to
367 exceed the concentration threshold C^* . When considering transient pumping, we note that the
368 mean contaminant concentration (Figure 4b) and the probability of exceedance (Figure 9)
369 display similar patterns for the lowest values of the variance of the flux-averaged concentration
370 observed at the pumping well (Figure 5b). This correspondence is not as evident for increased
371 concentration variance (Figure 5b). As noted in Section 4.2, highly variable transient pumping
372 strategies favor an increase of the uncertainty of solute concentrations at the pumping well.
373 This, in turn, affects risk uncertainty, as quantified by (5). In general, we observe that the
374 transient operational scheme selected in scenario III induces a temporal oscillating behavior of
375 risk at the well (Figure 9), according to which risk **tends** to decrease with the temporal
376 reduction of the mean concentration signal.

377 4.4. *Local scale dispersion and hydrogeological heterogeneity*

378 The previous sections considered a fixed level of heterogeneity and local scale dispersion (see
379 list of parameters in Table 1). We now assess the impact of aquifer heterogeneity and local
380 scale dispersion on the statistics of the flux-averaged concentration in the presence of variable
381 pumping rates.

382 4.4.1. *Impact of local scale dispersion on the concentration statistics*

383 Figure 10 illustrates the influence of increasing local scale dispersion on the mean and
384 on the variance of the flux-averaged concentration for the diverse pumping operations of
385 Scenario III. The mean concentration for constant and variable pumping rates are shown in
386 Figures 10a and 10b, respectively, whereas Figures 10c and 10d depict the concentration

387 variance for constant and variable pumping schedules, respectively. The effect of local scale
388 dispersion on the concentration variance is manifested through the longitudinal Péclet number,
389 defined as $P_e = \lambda v_0 / D_x$. Figure 10 shows the results for both longitudinal $P_e = 800$
390 (corresponding to the parameters listed in Table 1) and $P_e = 200$ for a fixed level of
391 heterogeneity (as rendered by $\sigma_Y^2 = 3$). The Péclet number of 200 corresponds to increased
392 local scale dispersion (longitudinal and transverse dispersivities are set to 0.04 m and 0.0004
393 m, respectively). The remaining input parameters for the simulations related to $P_e = 200$ are
394 the same as those reported in Table 1 of the manuscript.

395 As depicted in Figure 10, increasing the value of local scale dispersion (i.e., reducing P_e)
396 leads to decreased values of the concentration variance (Figures 10c and 10d). This is related
397 to the observation that dispersion smooths out the concentration gradients, hence reducing the
398 sample-to-sample variability of transport observables at the well (e.g., *Kapoor and Kitanidis,*
399 *1998; Fiori and Dagan, 2000; Dentz and de Barros, 2013*). Note that the transient operational
400 scheme applied at the well has marked influence on the concentration variance for both Péclet
401 numbers considered, as opposed to what can be observed for the constant in time pumping
402 scheme. This is mainly attributed to the following reasons. First, when the pumping rate is
403 temporally variable, the solute plume residence time in the aquifer increases, thus allowing for
404 the effects of local scale dispersion to become more pronounced. Second, temporally variable
405 flows, induced by the action of dynamic pumping rates, tend to enhance solute spreading,
406 which in turn augments dilution (see also *Dentz and Carrera, 2003*). Furthermore, this dilution
407 enhancement (driven by the action of variable pumping rates) leads to an increase of the plume
408 spatial extent (in the longitudinal and transverse directions) and to an increase in the probability
409 of the plume being captured by the well, thus contributing to a reduction of the variance. For

410 the set of parameters used in this numerical investigation, the significance of local scale
411 dispersion on the reduction of the concentration variance is diminished for a constant pumping
412 rate (compare Figure 10c and 10d). It is important to note that the joint effects of local scale
413 dispersion and pumping operations on the concentration variance will also depend on the
414 distance between the contaminant source and the receptor (i.e., the pumping well). With
415 reference to the effect of local scale dispersion on the concentration mean at the extracting well
416 (Figures 10a and 10b), we notice that the differences between the two values of P_e considered
417 are smaller as compared to the corresponding differences detected in the temporal evolution of
418 the concentration variance displayed in Figures 10c and 10d. When we increase local scale
419 dispersion, the mean concentration does not change much for a constant pumping schedule
420 (Figure 10a) because the distance between the source and the operating well is not sufficiently
421 large in our scenario (approximately equal to 10λ). Therefore, the effects of local scale
422 dispersion are not that evident on the temporal evolution of the mean concentration. We recall
423 that the advective time scale is the same for both values of P_e . On the other hand, for the case
424 of transient pumping (Figure 10b), the residence time of the plume in the aquifer increases,
425 thus allowing for the effects of P_e to become slightly visible. A further increase of the travel
426 distance of the contaminated plume will enhance the effects of local scale dispersion on the
427 mean concentration for both pumping operations (constant and variable in time).

428 4.4.2. [Impact of aquifer heterogeneity on the concentration statistics](#)

429 In the following, we investigate the sensitivity of the concentration statistics to the
430 degree of aquifer heterogeneity in the presence of both constant and time-varying pumping
431 schedule. The level of aquifer heterogeneity is epitomized by the log-conductivity variance σ_Y^2 .

432 Figure 11 provides a comparison between the mean (Figures 11a and 11b) and variance
433 (Figures 11c and 11d) of the flux-averaged concentration at the well for $\sigma_Y^2 = 0.5$ and 3 under
434 a constant and temporally dynamic pumping operations. The larger log-transmissivity variance
435 ($\sigma_Y^2 = 3$) yields an increased likelihood of the occurrence of preferential flow paths within the
436 aquifer which contribute to faster solute arrival. Solute migration through these high velocity
437 paths tends to be locally dominated by advection and the solute reaches the pumping well at
438 earlier times (Figures 11a and 11b). The temporal evolution of the mean concentration
439 displayed in Figures 11a and 11b for both log-transmissivity variances also shows that an
440 increased strength of σ_Y^2 can also augment the probability of occurrence of low conductivity
441 zones that can entrap solute, resulting in late arrival times. This contrast between low and high
442 transmissivity values for larger σ_Y^2 causes the plume to spread erratically and facilitates the
443 exchange of solute mass between streamtubes (i.e., dilution enhancement). In other words,
444 higher permeability variability tends to reduce the mean solute concentration and smooths out
445 concentration gradients, thus reducing concentration variance (Figures 11c and 11d). The
446 numerical results in Figures 11c and 11d illustrate that the differences between the
447 concentration variance obtained for $\sigma_Y^2 = 0.5$ and 3 are higher when the pumping operation
448 varies in time. The dynamic pumping scheme also increases the magnitude of the concentration
449 mean and variance when σ_Y^2 is lower (Figures 11b and 11d).

450 We remind that the results displayed in Figure 11 are not completely general because
451 other factors can play a role in the uncertainty of the concentration statistics. These additional
452 factors include the solute travel distance, source release conditions, location of the solute
453 source, the conceptualization of the heterogeneity model and the presence of chemical
454 reactions. Nevertheless, it is evident that the interplay between aquifer heterogeneity and

455 pumping operation plays fundamental role in probabilistic risk assessment. Figure 12 shows
456 the temporal variability of the risk (probability of exceedance), defined in Equation (5), for
457 constant and dynamic pumping operations and different values of σ_Y^2 . With the exception of
458 the early and late time regimes, the results in Figure 12 show that the probability of observing
459 a concentration value larger than a critical value is higher for low than for high heterogeneity.
460 The reason for this behavior is similar to the discussion related to Figure 11c and 11d; i.e.,
461 larger values of σ_Y^2 facilitate mass transfer between neighboring streamtubes, a phenomenon
462 which enhances dilution and, as a consequence, results in a larger probability of having lower
463 concentration values. Therefore, the effects of the interplay between advection and local-scale
464 dispersion become more evident for larger σ_Y^2 , resulting in a tendency to lower the risks at the
465 operating well (see Figure 12 for both pumping schemes). Otherwise, larger heterogeneity
466 leads to higher risks at the early and late times. We highlight that the risk computed for an
467 aquifer with high heterogeneity persists for longer periods of times (compare results obtained
468 for $\sigma_Y^2 = 0.5$ and $\sigma_Y^2 = 3$). This is a consequence of the increased tailing effects of the
469 concentration BTC observed in each MC realization when heterogeneity is large. For the
470 numerical set-up adopted in this work, we show that the pumping operation scheme has a
471 significant controlling role on the temporal evolution of the risk. As depicted in Figure 12b,
472 the temporal variability of the risk and its multiple peaks are controlled by the temporal patterns
473 of the pumping rate.

474 Albeit the focus of the current work lies on the interplay between aquifer heterogeneity
475 and pumping operation (i.e., engineering factors) on the uncertainty of the concentration BTC,
476 we include results in the SM for a homogeneous aquifer (i.e., constant transmissivity field).
477 Figure SM21 shows how the pumping operation can affect the asymmetry of the concentration

478 BTC under uniform flow fields. Furthermore, as shown in Figures SM20 and SM21, the peak
479 concentration in the deterministic homogeneous scenario is higher for both pumping operations
480 when compared to the heterogeneous aquifers. Figure 13 depicts the temporal evolution of the
481 concentration mean for different values of aquifer heterogeneity ($\sigma_Y^2 = 0.5$, $\sigma_Y^2 = 3$ and for the
482 homogeneous case). The increase of the mean concentration with lower heterogeneity is linked
483 to the selected location of the solute source zone (recall that, in the settings we analyze, the
484 centroid of the source zone is aligned with the location of the well). Plume meandering is less
485 pronounced for lower heterogeneous aquifers (when compared to higher heterogeneity in Y),
486 therefore the probability of the plume being captured by the well is higher for low values of
487 σ_Y^2 in our settings. Figures SM20 and SM21 depict the suite of results for the case of a
488 deterministically homogeneous aquifer ($\sigma_Y^2 = 0$).

489 **5. Conclusions**

490 We study and compare the effect of temporally variable and uniform pumping regimes
491 on key features of contaminant transport in a randomly heterogeneous aquifer. Our work
492 considers the joint effects of spatially heterogeneous hydraulic conductivity (or transmissivity)
493 and temporally varying well pumping rates and offers some insights on a realistic approach for
494 the evaluation of the risk associated with contamination of groundwater extracted at the
495 pumping well location. The analysis is performed within a stochastic framework upon relying
496 on the numerical study of three distinct pumping scenarios.

497 The two leading statistical moments of the flux-averaged contaminant concentrations
498 recovered at the well are computed. We document the way the temporally dynamic pumping
499 rate augments the uncertainty in the flux-averaged concentration at the well. We then provide

500 an appraisal of risk at the well by computing the probability that the contaminant concentration
501 exceeds a defined threshold value (i.e., probability of failure).

502 Our work leads to the following major conclusions.

503 1. In addition to aquifer heterogeneity, the use of a transient pumping strategy at the well can
504 markedly affect the temporal evolution of contaminant concentration BTCs and statistical
505 moments. We show that dynamic pumping strategies induce multiple peaks in the mean
506 and variance of the flux-averaged concentration detected at the extraction well.

507 2. Lowest and largest values of mean and variance of flux-averaged concentration at the well
508 tend to occur at the same time. This observation is supported by our numerical findings
509 and by analytical results illustrated in Appendix A.

510 3. The activation of the pumping well (or the increase of its extraction rate) induces
511 contaminant dilution with fresh water at the well. As a consequence, the detected
512 contaminant concentration tends to be reduced.

513 4. The choice of the type of engineering control to the temporal sequence of well pumping
514 rates could represent a key factor in quantifying the uncertainty of the contaminant
515 concentration detected at the well. This observation is supported by considering that
516 uncertainty associated with detected BTCs at the well increases for highly variable
517 pumping regimes (compare scenario III against scenarios I and II). We show that this
518 controlling role of the temporal extraction schedule on uncertainty has direct consequences
519 to risk analysis. The selection of a dynamic pumping regime has a clear influence on the
520 temporal evolution of risk, as defined in (5), at the well, i.e., pumping rate fluctuations
521 induce a temporally oscillating risk pattern. The mean flux-averaged solute concentration
522 and the probability of exceedance of a given threshold value show a similar temporal

523 evolution when the variance of flux-averaged concentrations at the pumping well is small.
524 In these cases, larger mean concentration values correspond to larger probability of
525 exceedance of a given concentration threshold. Otherwise, this correspondence was not
526 detected in our cases for increased concentration variance.

527 Our findings suggest that risk analysis should incorporate dynamic pumping rates to
528 reflect realistic operational practices and be able to provide important information to risk
529 managers, including realistic quantifications of the uncertainty associated with risk at the
530 pumping well.

531 Note that the results and conclusions presented in this study are confined to a
532 hydrogeological setting whose randomly heterogeneous conductivity is characterized by given
533 sets of parameters. The results can be impacted by other factors, including, e.g., the distance
534 between the contaminant source and the receptor, the dimensions of the source zone, mass
535 release conditions, the ratio of the base flow discharge with respect to the pumping rate,
536 chemical reactions and the conceptualization of the hydrogeological heterogeneity. The
537 influence of variable pumping rates on probabilistic risk analysis depends on the interactions
538 between temporal fluctuations induced by the action of pumping, local scale dispersion, and
539 spatial heterogeneity. Quantifying the effects of the overall spectrum of scenarios is complex
540 due to the interplay of three distinct time scales defined by the pumping well operation,
541 advection and local scale dispersion. The relative importance of the temporally variable
542 pumping rate on risk will depend on the ratio between these characteristic time scales and
543 should be subject of further investigation. As a future projection of the study, one can
544 incorporate in our methodological framework adverse human health effects (e.g., *Andričević*
545 *and Cvetković*, 1996, *de Barros and Rubin*, 2008) and other well vulnerability criteria,

546 including, e.g., the time taken to breach a certain quality objective (e.g., drinking standard) at
547 the well, the total time of well failure (i.e., non-compliance with a quality objective), and the
548 time taken to recover the well from failure.

549

550 **Appendix A: Low-order moments of the flux-averaged concentration**

551

552 This appendix provides theoretical developments supporting the observed
553 correspondence between the low and high values in the mean and variance of the flux-averaged
554 concentration (see Figures 4-5). We start by considering the migration of a solute plume
555 originating far from the operating pumping well from a source zone of volume V_0 located
556 upstream of the well. To simplify the derivation, we quantify the mean and variance of mass
557 flux at a control plane located far away from the well and spanning the wellhead protection
558 area (WHPA) along a direction normal to the mean background groundwater flow (see Figure
559 A.1). Under this setting, we assume that the pumping rate at the well is constant such
560 that $Q_w(t) = Q_w$ and develop theoretical expressions for the moments of flux-averaged
561 concentrations under a uniform-in-the-mean flow condition. As such, the key purpose of our
562 subsequent developments is to establish an analogy with the observations stemming from the
563 numerical results illustrated in Section 4. We obtain an approximation for the mean and
564 variance of the flux-averaged concentration at the control plane by making use of the
565 Lagrangian framework developed by *Dagan et al.* (1992). For further details on the Lagrangian
566 approach, the reader is referred to Chapters 9 and 10 of *Rubin* (2003). Given the purpose of

567 the analogy, we neglect the effects of local-scale dispersion (e.g., transport is purely advective)
 568 in the following derivations and discussion.

569 We start by defining $C(t)$ as the flux-averaged contaminant concentration [M/L³]
 570 measured at the control plane described above (denoted by CP)

571

$$C(t) = \frac{Q_m(t)}{Q_w}. \quad (\text{A.1})$$

572

573 We respectively denote $Q_m(t)$ [M/T] and Q_w [L³/T] as the solute mass and volumetric flow
 574 rate across the CP at time t . Since Q_w is deterministic in our work, the statistics of $C(t)$ depend
 575 only on the statistics of $Q_m(t)$. For initial concentration c_o instantaneously and uniformly
 576 injected within a volume V_0 , the mean of $Q_m(t)$ can be expressed as:

577

$$\langle Q_m(t|\tilde{\mathbf{a}}) \rangle = C_0 \int_{V_0} g_1(t|\tilde{\mathbf{a}}) d\tilde{\mathbf{a}}. \quad (\text{A.2})$$

578

579 with $\tilde{\mathbf{a}}$ indicating the initial location (within V_0) of a solute particle released in the aquifer,
 580 $g_1(t|\tilde{\mathbf{a}})$ being the solute travel time probability density function (PDF) from the source to the
 581 CP. The second moment of $Q_m(t)$ can be expressed as:

582

$$\langle [Q_m(t|\tilde{\mathbf{a}})]^2 \rangle = C_0^2 \int_{V_0} \int_{V_0} g_2(t, t|\tilde{\mathbf{a}}', \tilde{\mathbf{a}}'') d\tilde{\mathbf{a}}' d\tilde{\mathbf{a}}'', \quad (\text{A.3})$$

583

584 where $g_2(t, t|\tilde{\mathbf{a}}', \tilde{\mathbf{a}}'')$ is the two-particle travel time PDF. The variance of $Q_m(t)$ is obtained
 585 by evaluating $Var[Q_m] = \langle [Q_m]^2 \rangle - \langle Q_m \rangle^2$. For low heterogeneity, e.g. $\sigma_Y^2 < 1$, the PDF g_1
 586 can be represented by a lognormal distribution (*Dagan et al.*, 1992; *Cvetkovic et al.*, 1992;
 587 *Selroos and Cvetkovic*, 1994). This result has been verified numerically (*Rubin*, 2003; *Gotovac*
 588 *et al.*, 2009). Under this assumption, g_1 scales as:

$$g_1(t) \sim \frac{1}{t} e^{-(\ln t)^2}. \quad (\text{A.4})$$

590

591 By the same token, the two-particle travel time PDF g_2 scales as:

$$g_2(t) \sim \frac{1}{t^2} e^{-(\ln t)^2}. \quad (\text{A.5})$$

592

593 We observe that $\sqrt{g_2}$ evolves like g_1 . Thus, the temporal evolution of the mean and
 594 variance of $Q_m(t)$ (and consequently of $C(t)$) are similar (see (A.2) and (A.3)). This
 595 theoretical finding, albeit under the simplified conditions here considered, constitutes an
 596 additional support to the temporal coincidence between maxima and minima of the mean and
 597 the variance of $C(t)$ observed in our work. Note that the travel time scaling we derive is strictly
 598 valid for low values of log-conductivity/transmissivity variance. It can nevertheless serve as
 599 an approximation for moderate to high variances as seen for example in the work of *Salandin*
 600 *and Fiorotto* (1998) who found a good quality agreement between longitudinal velocity
 601 covariances obtained through numerical Monte Carlo simulations and first-order theory for
 602 log-conductivity variance as large as 4 under uniform-in-the-mean flow.

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611

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751

(L_x, L_y)	Flow Domain	(170 m, 150 m)
(Δ_x, Δ_y)	Grid size	$1/8 \lambda$ m, $1/8 \lambda$ m
λ	Correlation length of Y	8 m
σ_Y^2	Variance of Y	3
ϕ	Porosity	0.2
J	Mean head gradient	0.005882
T	Average Transmissivity	$1 \text{ m}^2/\text{d}$
D_x	Longitudinal dispersion coefficient	$2.94 \times 10^{-4} \text{ m}^2/\text{d}$
D_y	Transverse dispersion coefficient	$2.94 \times 10^{-6} \text{ m}^2/\text{d}$
(x_w, y_w)	Location of the pumping well	(114.5 m, 74.5 m)
Q_w	Temporal sequence of pumping rate for scenario I	0.6 (uniform pumping rate) vs 0.8/0.4/0.8/0.4/0.8/0.4/0.8/0.4/0.8/0.4/0.8/0.4/0.8/0.4/0.8/0.4 m^3/d
Q_w	Temporal sequence of pumping rate for scenario II	0.45 (uniform pumping rate) vs 0.8/0.7/0.6/0.5/0.4/0.3/0.2/0.1/0.8/0.7/0.6/0.5/0.4/0.3/0.2/0.1 m^3/d
Q_w	Temporal sequence of pumping rate for scenario III	0.3 (uniform pumping rate) vs 0.8/0/0.4/0/0.8/0/0.4/0/0.8/0/0.4/0/0.8/0/0.4/0 m^3/d
q_o	Groundwater specific discharge	$5.882 \times 10^{-3} \text{ m}/\text{d}$
V_w	Volume of extracted water for scenario I	2880 m^3
V_w	Volume of extracted water for scenario II	2160 m^3
V_w	Volume of extracted water for scenario III	1440 m^3
C^*	Contaminant concentration threshold	$10 \text{ g}/\text{m}^3$

Table 1. Parameter set employed in the numerical simulations.

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Figure Captions

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Fig. 1. Schematic representation of the problem studied. A two-dimensional flow field characterized by the presence of a pumping well and a uniform (in the mean) base flow from left to right of the domain is considered. The contaminant source is rectangular with longitudinal dimension L_1 and transverse dimension L_2 , located at distance L from the well positioned at $\mathbf{x}_w = (x_w, y_w)$. The contaminant source is aligned with respect to the pumping well center.

Fig. 2. Pumping scenarios analyzed in the study: Scenario (a) I; (b) II; (c) III.

Fig. 3. Contaminant concentration BTCs observed at the pumping well for the three scenarios analyzed: Scenario I (plots a, b, c); Scenario II (plots d, e, f); and Scenario III (plots g, h, i).

Fig. 4. Mean concentration $\langle C \rangle$ observed at the pumping well for the uniform pumping strategy of Scenario III (a) and the transient pumping strategy of Scenario III (b).

Fig. 5. Variance of the concentration $Var[C]$ observed at the pumping well for the uniform pumping strategy of Scenario III (a) and the transient pumping strategy of Scenario III (b).

774 **Fig. 6.** Coefficient of variation CV of the concentration observed at the pumping well for the
775 uniform pumping strategy of Scenario III.

776

777 **Fig. 7.** Coefficient of variation CV of the concentration observed at the pumping well for the
778 transient pumping strategy of Scenario III.

779

780 **Fig. 8.** Sample probability of exceedance of the concentration threshold $P(C > C^*)$ for the
781 uniform pumping strategy of Scenario III.

782

783 **Fig. 9.** Sample probability of exceedance of the concentration threshold $P(C > C^*)$ for the
784 transient pumping strategy of Scenario III.

785

786 **Fig. 10.** Impact of the Péclet number (P_e) on the concentration mean for (a) constant and (b)
787 time-varying pumping rate. Impact of the Péclet number (P_e) on the concentration variance
788 for (c) constant and (d) time-varying pumping rate. Purple curves refer to $P_e = 800$ and blue
789 curves refer to $P_e = 200$.

790

791 **Fig. 11.** Mean concentration for (a) constant and (b) time-varying pumping rates. Variance
792 behavior is also depicted for (c) constant and (d) time-varying pumping schemes. Blue curves
793 refer to low heterogeneity and purple curves to high heterogeneity.

794

795 **Fig. 12.** Probability of exceedance of the concentration threshold $P(C > C^*)$ for (a) constant
796 and (b) time-varying pumping rates. Blue curves refer to low heterogeneity and purple curves
797 to high heterogeneity.

798

799 **Fig. 13.** Mean concentration $\langle C \rangle$ observed at the pumping well for (a) constant and (b) time-
800 varying pumping strategy of Scenario III for different levels of aquifer heterogeneity.

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802 **Fig. A.1.** Schematic representation of the problem analyzed using the Lagrangian framework.

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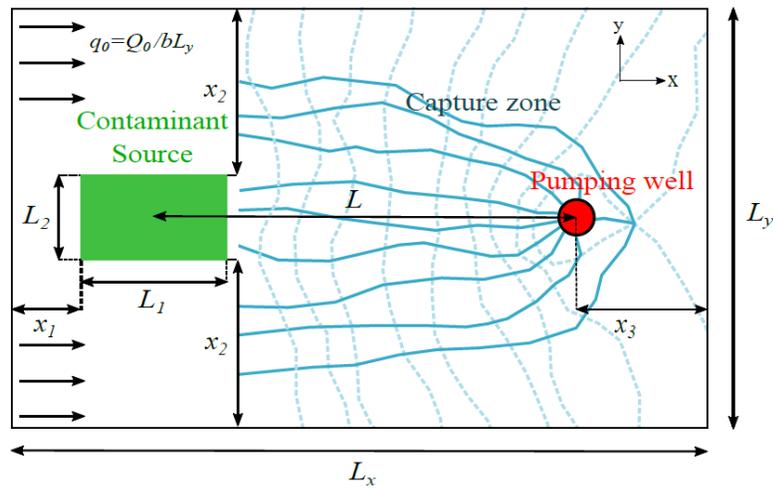
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Figures and Captions



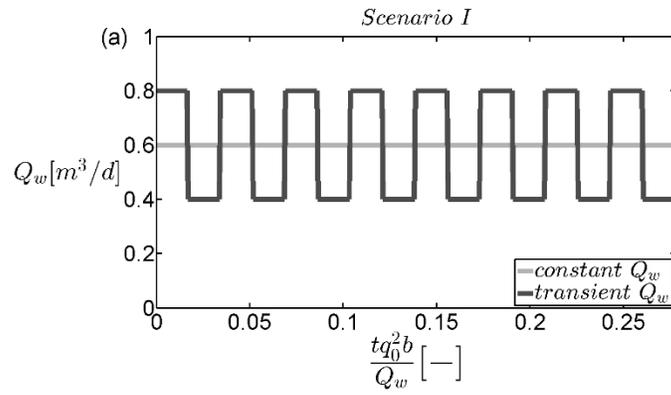
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819 well center.

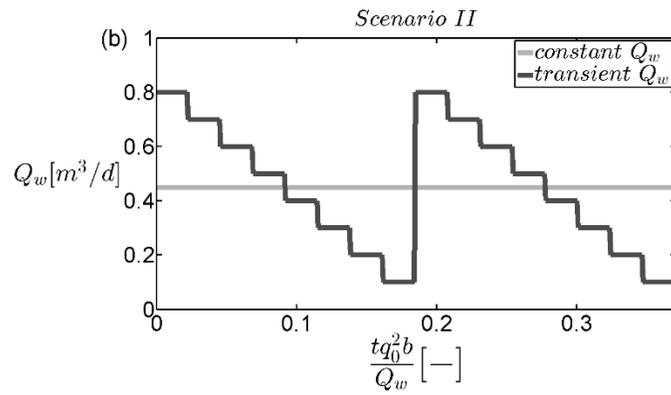
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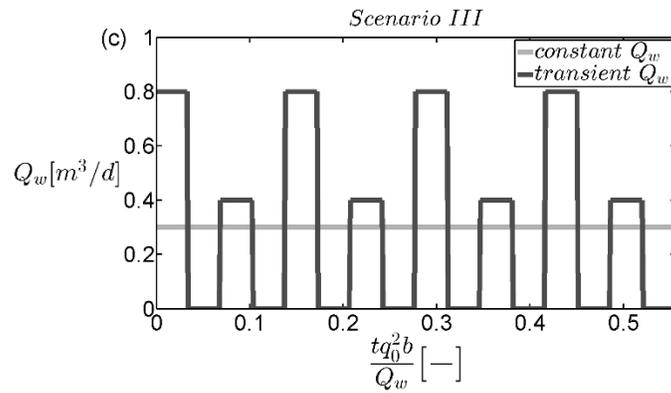
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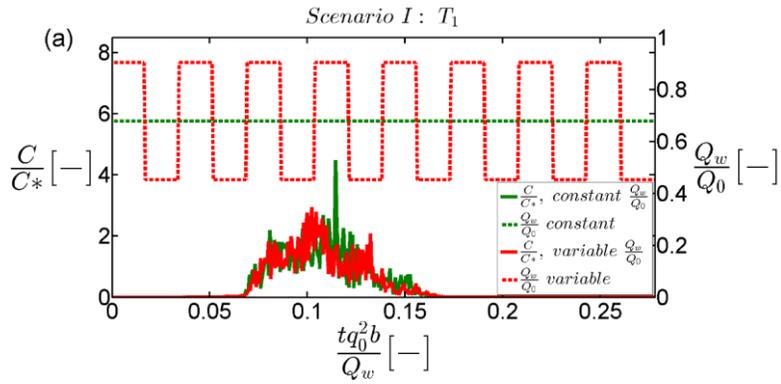
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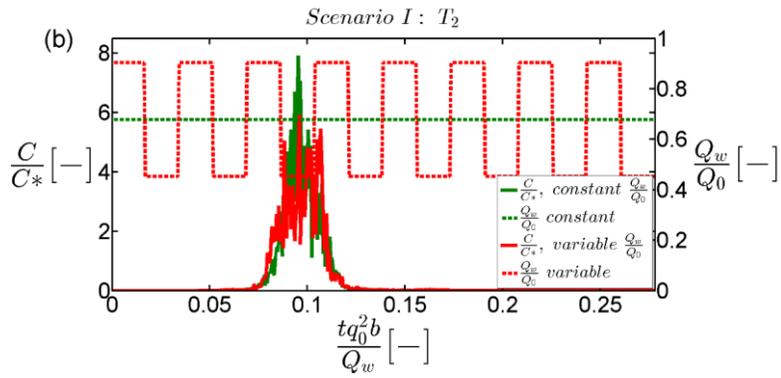
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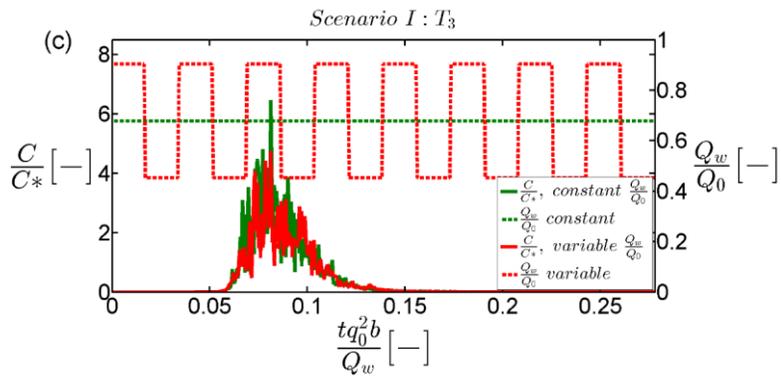
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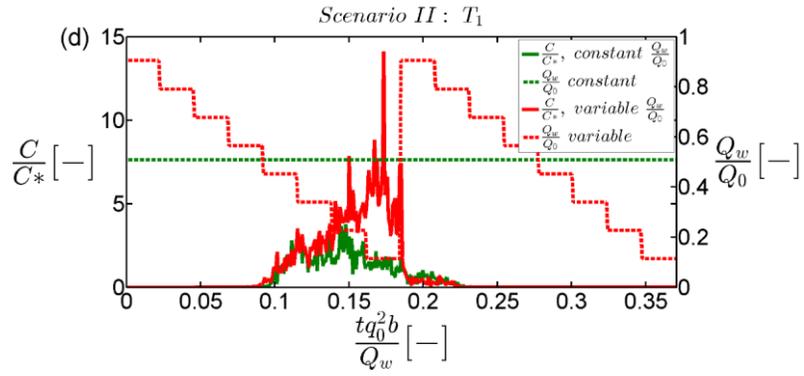


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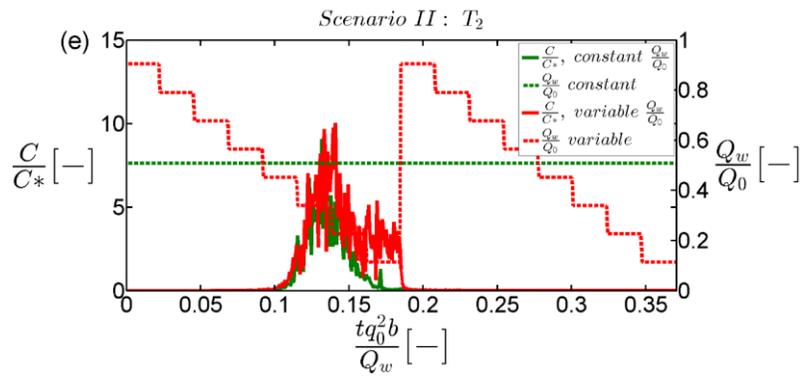


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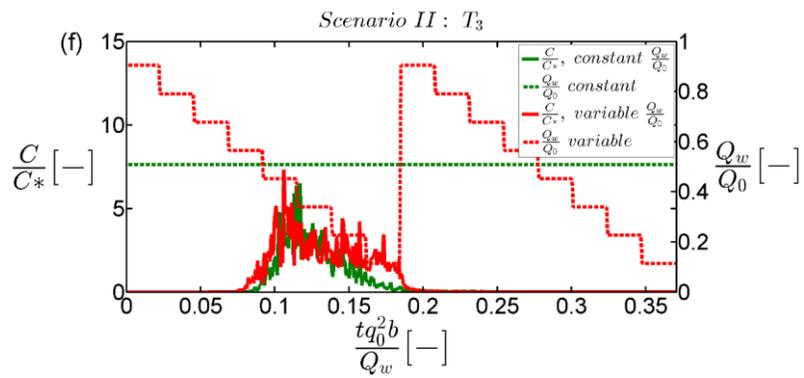




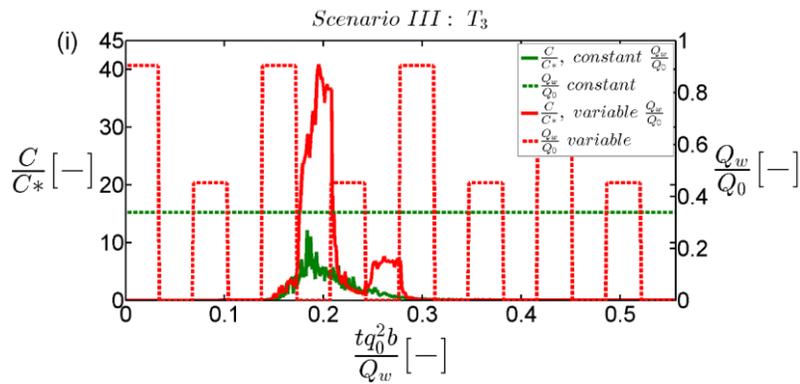
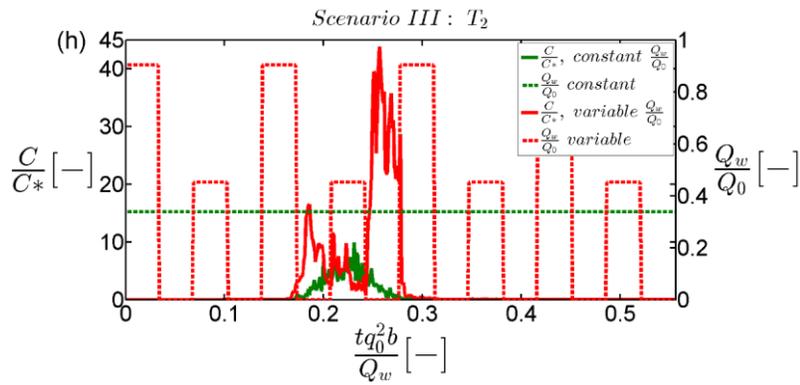
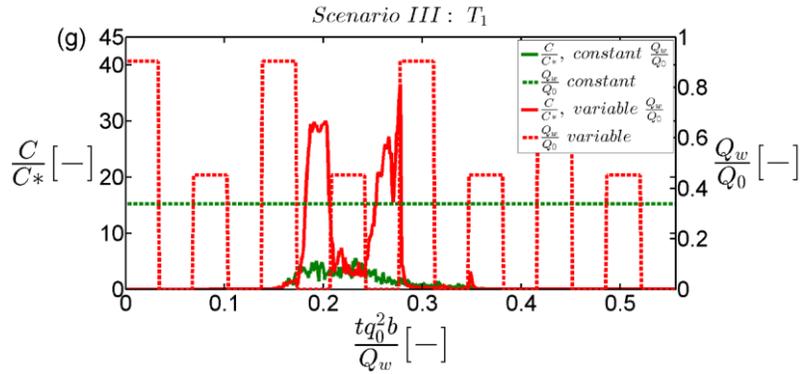
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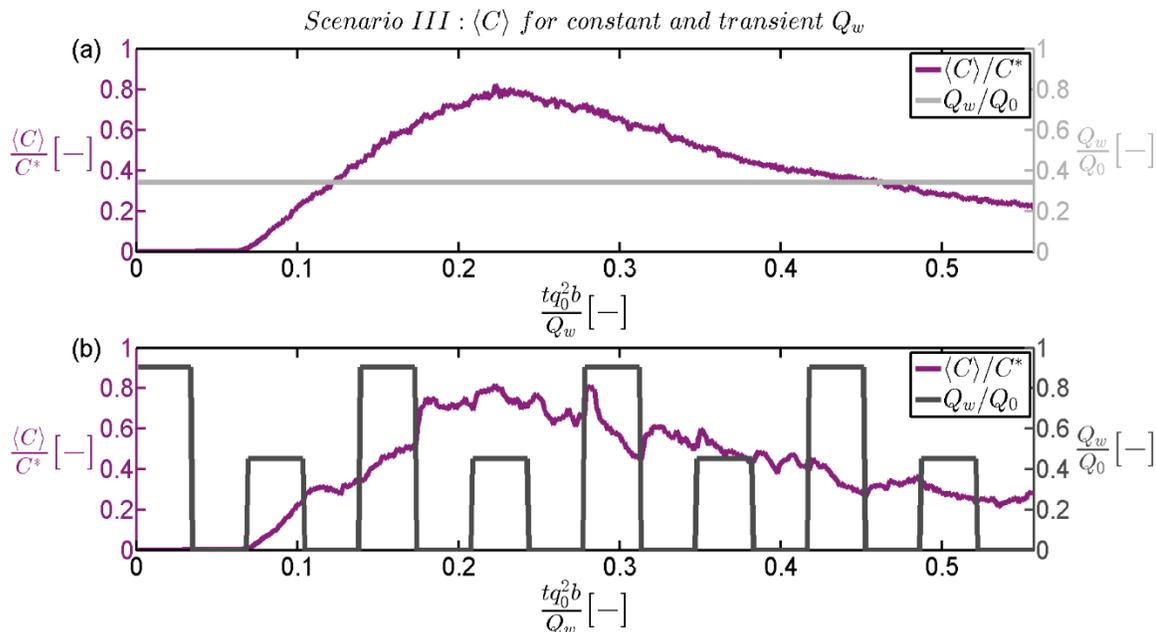
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837 **Fig. 3.** Contaminant concentration BTCs observed at the pumping well for the three
 838 scenarios analyzed: Scenario I (plots a, b, c); Scenario II (plots d, e, f); and Scenario
 839 III (plots g, h, i).

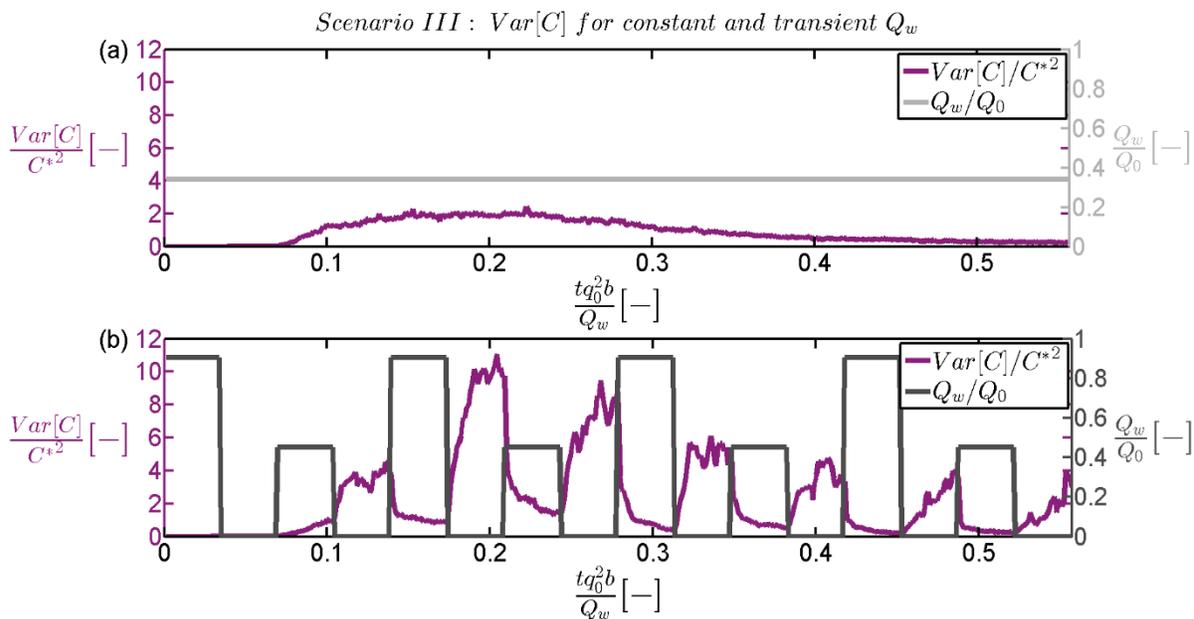
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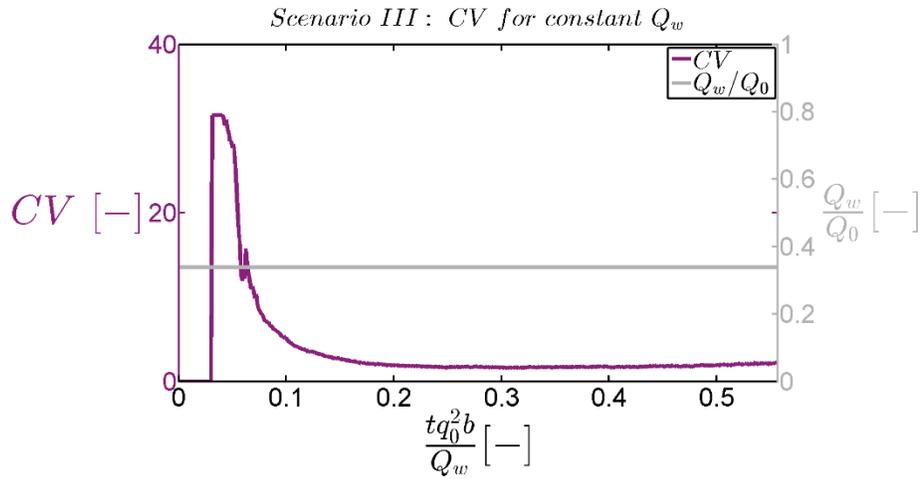
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843 **Fig. 4.** Mean concentration $\langle C \rangle$ observed at the pumping well for the uniform pumping
 844 strategy of Scenario III (a) and the transient pumping strategy of Scenario III (b).



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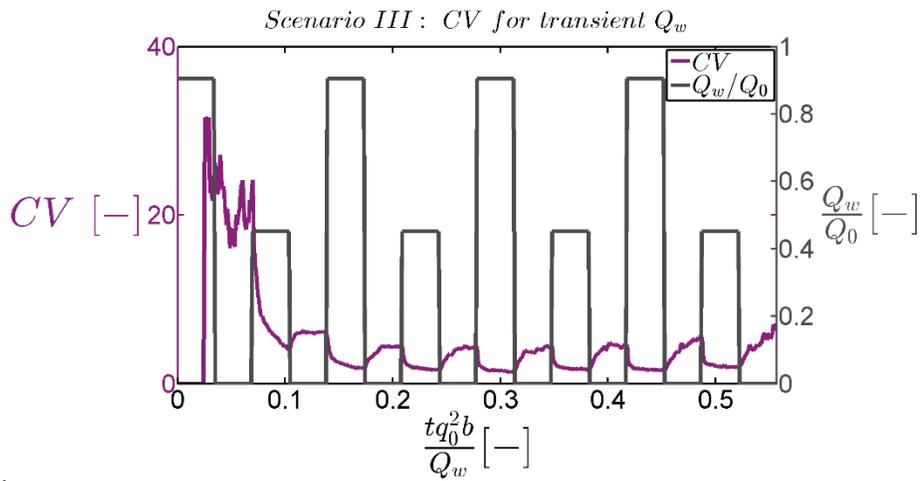
846 **Fig. 5.** Variance of the concentration $Var[C]$ observed at the pumping well for the uniform
 847 pumping strategy of Scenario III (a) and the transient pumping strategy of Scenario III (b).



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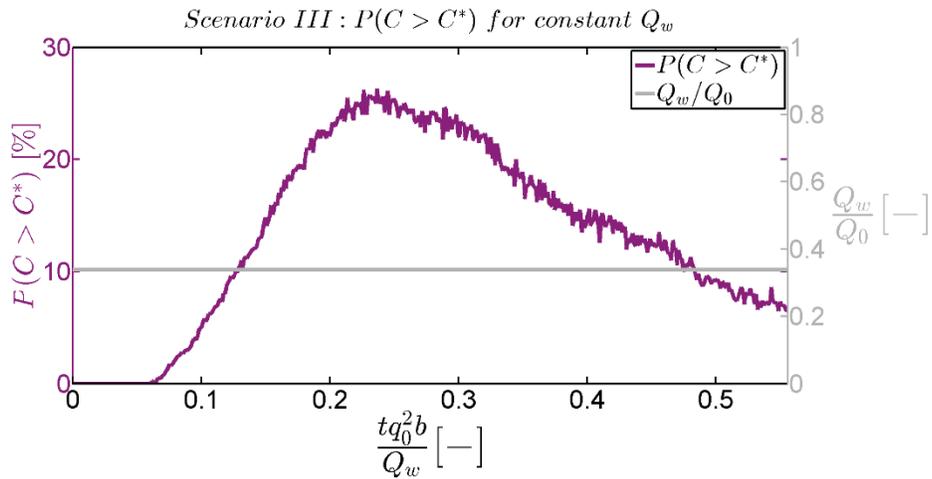
849 **Fig. 6.** Coefficient of variation CV of the concentration observed at the pumping well for the
 850 uniform pumping strategy of Scenario III.

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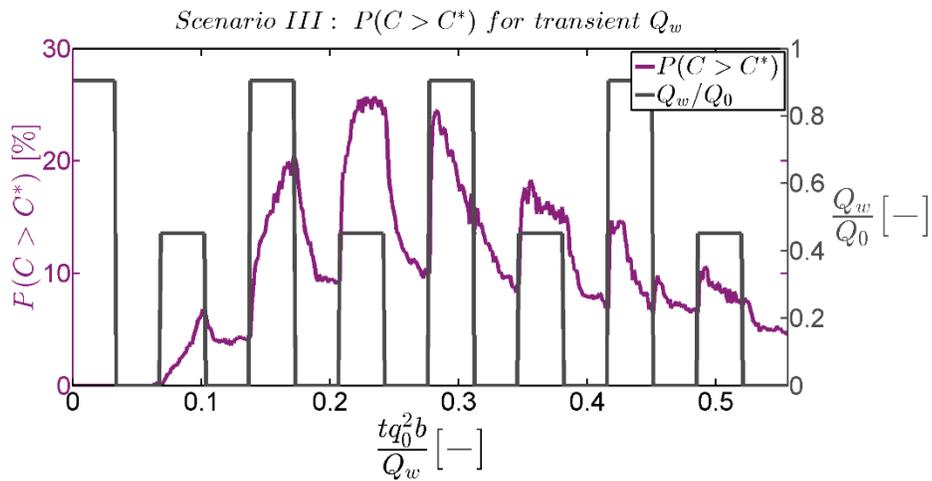
853 **Fig. 7.** Coefficient of variation CV of the concentration observed at the pumping well for the
 854 transient pumping strategy of Scenario III.



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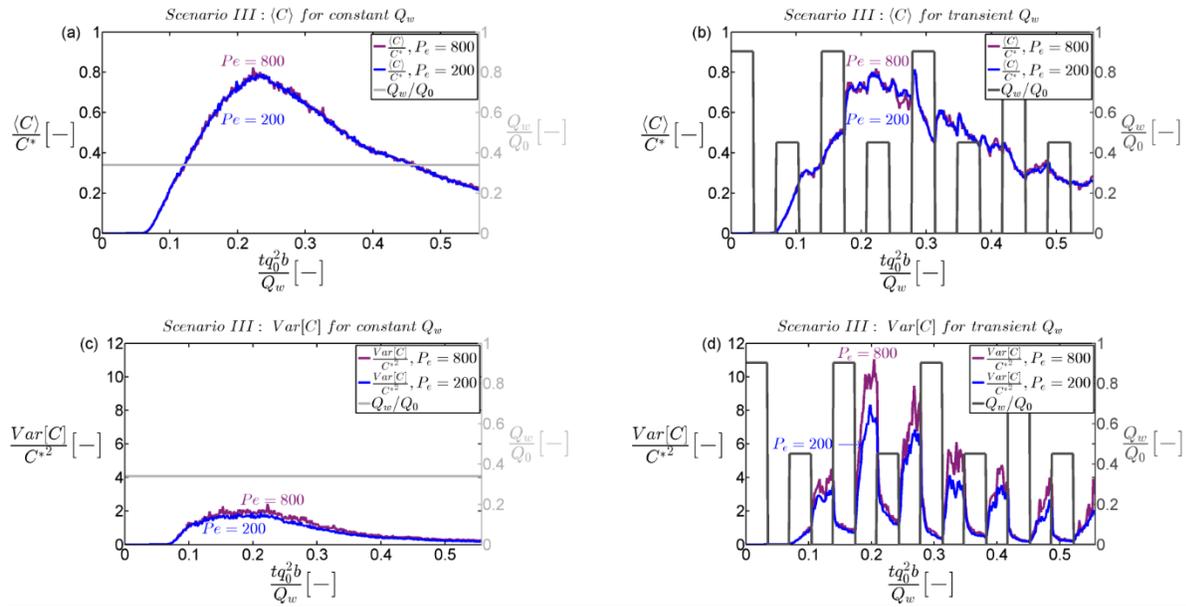
856 **Fig. 8.** Probability of exceedance of the concentration threshold $P(C > C^*)$ for the uniform
 857 pumping strategy of Scenario III.

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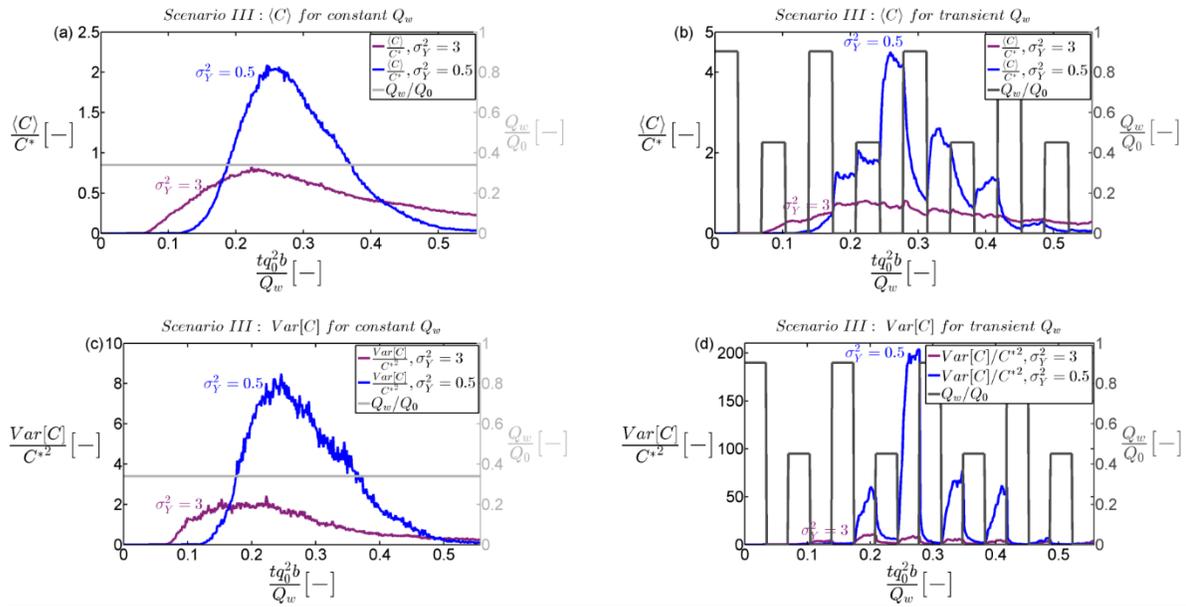
860 **Fig. 9.** Probability of exceedance of the concentration threshold $P(C > C^*)$ for the transient
 861 pumping strategy of Scenario III.



862

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 864 time-varying pumping rate. Impact of the Péclet number (P_e) on the concentration variance
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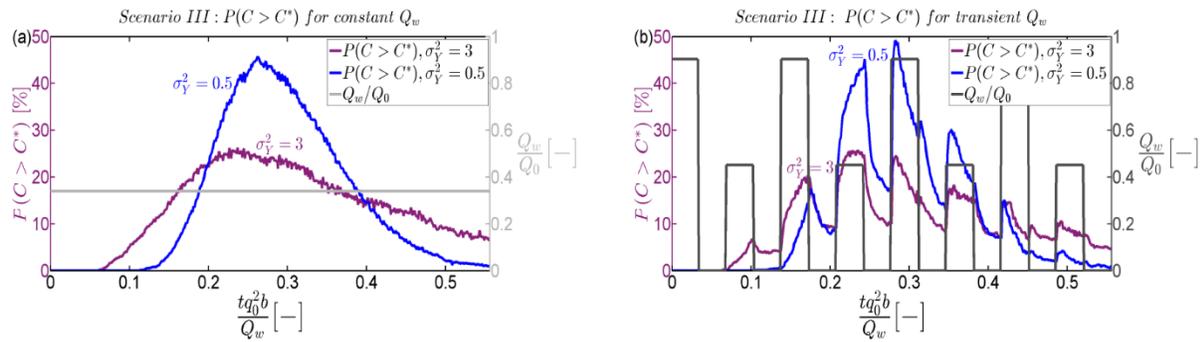
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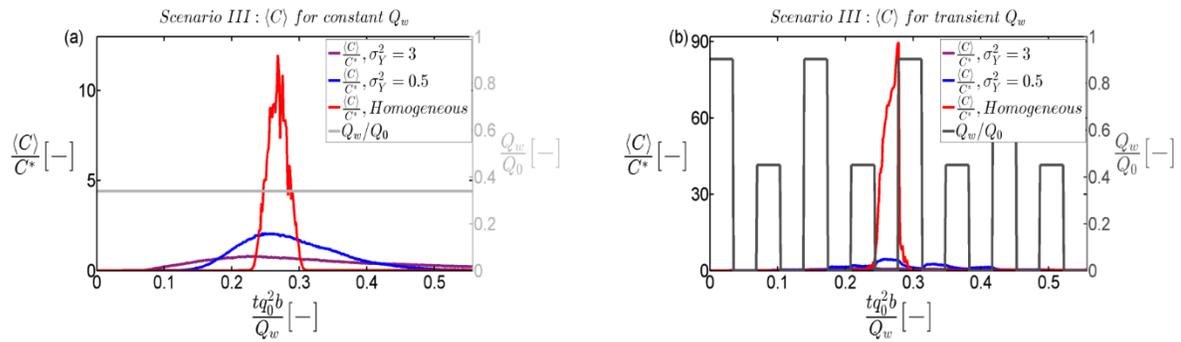


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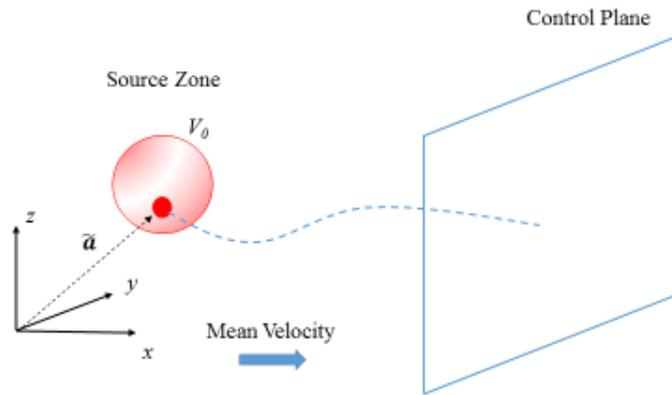
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880 **Fig. 13.** Mean concentration $\langle C \rangle$ observed at the pumping well for (a) constant and (b) time-
 881 varying pumping strategy of Scenario III for different levels of aquifer heterogeneity.

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883

884 **Fig. A.1.** Schematic representation of the problem analyzed using the Lagrangian framework.

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