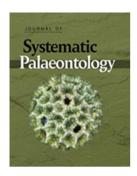
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A silicified Early Triassic marine assemblage from Svalbard

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- 1 A silicified Early Triassic marine assemblage from Svalbard
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Abstract

Understanding how the marine biosphere recovered from the late Permian mass extinction event is a major evolutionary question. The quality of the global fossil record of this interval is, however, somewhat poor due to preservational, collection and sampling biases. Here we report a new earliest Induan (Hindeodus parvus Zone) marine assemblage from the Deltadalen Member of the Vikinghøgda Formation, central Spitsbergen, which fills a critical gap in knowledge. The fully silicified fossils comprise the oldest silicified assemblage known from the Triassic and provide critical new systematic data. For its age, the assemblage is exceptionally diverse with 14 species of bivalves and gastropods, as well as conodonts and ammonoids. Four new bivalve species (Austrotindaria antiqua, A. svalbardensis, Nucinella taylori and N. nakremi) and one new gastropod species (Glabrocingulum parvum) are described, and five families are recorded in the Induan for the first time. Some of the common and globally widespread Early Triassic taxa, such as *Unionites*, are also present and their exceptional preservation reveals key morphological characters that are documented for the first time. Taxonomic and ecological revisions based on these new data suggest that shallow-infaunal deposit-feeders were a dominant component of pre-Spathian benthic communities. The gastropods and bivalves all possessed a planktotrophic larval stage, which may have been a particular advantage in the wake of the late Permian mass extinction.

- **Keywords**: extinction; Early Triassic; diversification; Mollusca; Protobranchia;
- 31 chemosymbiosis.

Introduction

The aftermath of the late Permian mass extinction represents a key interval in the diversification of marine biota. Even though 78% of marine genera are estimated to have gone extinct during the late Permian event no novel phyla or classes and only one new mode of life originated during the extinction aftermath (Erwin et al. 1987; Foster & Twitchett 2014). The re-diversification of benthic groups in the wake of the extinction, e.g. bivalves (Nakazawa & Runnegar 1973; Komatsu et al. 2004; Fraiser & Bottjer 2005; Hautmann 2007; Posenato 2008) and gastropods (Erwin & Pan 1996; Nützel 2005; Gründel & Nützel 2012), is typically described as occurring in the Middle Triassic, resulting in the traditional view of a 'delayed' post-Permian recovery. The Early Triassic fossil record is, however, notoriously poor, as evidenced by an unusually high number of Lazarus taxa, due to the typical mouldic preservation of shells (Nakazawa & Runnegar 1973; Wheeley & Twitchett 2005), their small, easily overlooked, sizes (Hautmann & Nützel 2005), as well as the substantial regional and latitudinal sampling bias towards low (tropical) palaeolatitudes of Palaeotethys (Foster & Twitchett 2014).

Better preserved benthic fossil assemblages have been reported from the Induan of South China (Kaim et al. 2010; Hautmann et al. 2011; 2015) and Primorye, Russia (Shigeta et al. 2009), and from the Olenekian of the western USA (Batten & Stokes 1987; Hautmann & Nützel 2005; Nützel & Schulbert 2005; Pruss et al. 2015) and Pakistan (Wasmer et al. 2012; Kaim et al. 2013). In addition, a partially silicified fauna is known from the Griesbachian of Oman (Twitchett et al. 2004; Wheeley & Twitchett 2005; Oji & Twitchett 2015). Studies of those assemblages have shown that a greater number of bivalve, gastropod and crinoid lineages survived the late Permian mass extinction event and/or were present in the Early Triassic than previously thought. These partially silicified and better preserved assemblages still lack key diagnostic characters of many taxa, however, especially the internal morphology of bivalve shells, making taxonomic assignments often equivocal.

Better preserved early Induan (Griesbachian) fossil assemblages from palaeotropical localities have been critical in demonstrating that taxonomically and ecologically diverse benthic ecosystems reappeared locally in some settings by the second conodont zone of the Triassic in Neotethys (Twitchett et al. 2004). At higher latitudes, however, early Induan shelly benthic assemblages are rarely reported and typically consist of poorly preserved cosmopolitan generalists, such as *Claraia, Unionites, Lingularia* and *Warthia* (Spath 1930; Mørk et al. 1999; Zonneveld et al. 2007). In these regions, trace fossils are typically used as a

measure of documenting ecosystem collapse and recovery (following Twitchett 2006) and demonstrate that rapid local recovery of the infauna took place within the earliest Induan (Twitchett & Barras 2004; Beatty et al. 2008; Zonneveld et al. 2010). From their analysis of global fossil occurrences, Foster & Twitchett (2014) also concluded that during the Induan the benthic fauna of the extratropical northern palaeolatitudes contained a greater ecological diversity than that of the palaeotropics.

Here, we describe the first silicified fossil assemblage from the Lower Triassic of the northern palaeolatitudes. The assemblage was recovered from the lower Griesbachian of central Spitsbergen, Svalbard, making it the oldest silicified assemblage known from the Mesozoic. It is the first fully silicified assemblage of Early Triassic age and the exquisite preservation of the internal and external morphology of the bivalve and gastropod shells provides critical new taxonomic data that have major implications for our understanding of the response of marine ecosystems to the late Permian mass extinction event and of the stratigraphic ranges of several molluscan families.

Materials

The fossil assemblage described herein was recovered from two carbonate concretions, collected by WJF & RJT in 2013 from 11.9 and 12.6 m above the base of the Deltadalen Member of the Vikinghøgda Formation, Lusitaniadalen, Svalbard (N78° 17′ 54.8″, E016° 43′ 59.3″; Fig. 1). The fossils were extracted in the laboratory by first mechanically disaggregating 3kg samples to expose larger fossils. Remaining rock fragments were then dissolved using the buffered formic acid technique of Jeppsson & Anehus (1995). To maximise yield, the residue was collected at ca. 12-hour intervals, washed thoroughly with tap water to remove any excess solution and to avoid crystal growth, and dried. The buffered solution was renewed every 48 hours. Heavy liquid separation (Mitchell & Heckert 2010) was used to separate the fossils from the remaining residue. The specimens are housed in the Natural History Museum, London.

Geological setting

During the Early Triassic, the Svalbard archipelago was situated at ca. 45-50°N in the southern part of the Boreal Ocean (Fig. 1; Hounslow et al. 2008). In central Spitsbergen, the Vikinghøgda Formation, described by Mørk et al. (1999), records deposition through the

latest Permian and entire Early Triassic in a siliciclastic, open-marine, shelf setting. It is divided into three members, of which the Deltadalen Member is the lowest (Mørk et al. 1999). At the study site in Lusitaniadalen, the basal 1.6 m of the Deltadalen Member is composed of bedded, well-bioturbated, fine- to medium-grained, glauconitic sandstones that are very similar to the underlying Kapp Starostin Formation except that they lack diagenetic chert nodules (Mørk et al. 1999; Nabbefeld et al. 2010). These sandstones contain a diverse trace fossil assemblage, indicating a fully functional benthic ecosystem that was living under well-oxygenated conditions (Nabbefeld et al. 2010). In contrast, the body fossil assemblage is limited, comprising mainly the phosphatic-shelled lingulid brachiopods, which may indicate preservational bias.

The sandstones record the onset of marine transgression (Mørk et al. 1999; Nabbefeld et al. 2010) with the base of the overlying laminated, silty mudstones marking significant deepening. Several horizons of cemented siltstones and tabular concretions are found in the lower few metres of this mudstone-dominated interval and contain coarser laminae with abundant ichthyoliths and disarticulated lingulid shells. Occasional, 1-5cm thick, finegrained, pyritic and glauconitic, graded, cemented sandstones, interpreted as distal tempestites, also occur within the laminated silty mudstones. Biomarker data support the interpretation based on field observations that deposition took place under anoxic and periodically euxinic conditions during transgression and maximum flooding, with evidence of phytoplankton blooms in the surface waters likely driven by nutrient influx (Nabbefeld et al. 2010).

From ca. 10m above the base of the mudstones, the lithologies become noticeably coarser with a greater proportion of interbedded siltstones and very fine sandstones, presumably due to progradation or sealevel fall. The coarser, heterolithic beds are bioturbated, initially by mm-diameter *Planolites* and then with the addition of *Skolithos* and *Arenicolites* ca. 1.5 m higher, indicative of environmental amelioration and deposition under more persistently oxygenated conditions. Ichnofabric index remains low (ii2-3), however, and the burrows are small (diameters <5mm) and do not penetrate deeply, indicating that the environment was probably still not fully optimal for benthic colonisation. Earliest Triassic ammonoids, conodonts and a moderate benthic assemblage have already been recorded from the *Planolites*-dominated interval (Mørk et al. 1999). One of the concretions that yield the silicified assemblage documented in this study derives from that same horizon, 12.6 m above the base of the formation. The second concretion was collected from ca. 65 cm lower (Fig. 1).

The age of the Deltadalen Member is well constrained by biostratigraphy (ammonoids and conodonts) and magnetostratigraphy, with deposition recorded from the upper Changhsingian, prior to the late Permian mass extinction event, through most of the Induan (Mørk et al. 1999; Hounslow et al. 2008; Nakrem et al. 2008). Locally, the late Permian extinction event is recorded by the disappearance of prolific bioturbation just below the top of the glauconitic sandstones, 1.6m above the base of the member (Nabbefeld et al. 2010). From conodont evidence, Nakrem et al. (2008) inferred that the Permian/Triassic boundary occurs between 5 m and 11 m above the base of the Deltadalen Member. *Claraia* cf. *wangi* is recorded 10.1 m above the base of the member at Lusitaniadalen, indicative of a Griesbachian age. The silicified assemblages described herein are of earliest Griesbachian (earliest Induan) age, and are assigned to the upper part of the *Otoceras boreale* Zone, which corresponds to the basal Triassic *Hindeodus parvus* Conodont Zone (Orchard 2007).

Systematic palaeontology

Phylum **Brachiopoda** Duméril, 1805 Class **Lingulata** Goryansky & Popov, 1985 Order **Lingulida** Waagen, 1885 Family: **Discinidae** Gray, 1840 Genus *Orbiculoidea* d'Orbigny, 1847

Type species. Orbicula forbesii Davidson, 1848; Silurian, Wenlock; West Midlands, England.

Diagnosis. The generic diagnosis follows Mergl (2006). Shell thin, strongly dorsibiconvex, subcircular, with subtrapezoidal outline of the posterior part. Ornament on both valves composed of regular raised concentric fila separated by broader interspaces. Dorsal valve is conical to subconical with subcentral apex; ventral valve with subcentral apex. Ventral valve depressed conical with subcentral apex. Pedicle track narrow, tapering posteriorly and anteriorly closed by shallow listrium; foramen in posterior end of listrium with short internal tube.

Species *Orbiculoidea winsnesi* Gobbet, 1963 (Fig. 2A)

1963 Orbiculoidea winsnesi Gobbett, p. 46; Plate 1; Figs 4-5.

165	Material. One dorsal valve from LD-04 (<u>NHMUK PI BE 3238</u>).				
166	Description. Shell outline is sub-circular, with its length slightly greater than the width. The				
167	dorsal valve is inflated to an apex, which is located about a third of the diameter from the				
168	anterior margin. Shell smooth, except for fine concentric growth lines.				
169	Remarks. Extant solitary discinids are sessile, epifaunal, invertebrates that attach to hard				
170	surfaces with a sucker-like pedicle (Mergl 2010). They have been attaching to shelled				
171	invertebrates since the Ordovician (Mergl 2010), and one of the specimens described herein				
172	is attached to an ammonoid. Discinids are suspension feeders and their co-occurrence with				
173	lingulids in laminated black shales has led some authors to interpret them as being tolerant of				
174	low-oxygen conditions (Savoy 1992; Hallam 1995; Mergl 2010).				
175	Mode of Life. Surficial, stationary, attached, suspension feeder (Mergl 2010).				
176					
177	Phylum Mollusca Linnaeus, 1758				
178	Class Bivalvia Linnaeus, 1758				
179	Order Pterioida Newell, 1965				
180	Family Myalinidae Frech, 1891				
181	Genus <i>Promyalina</i> Kittl, 1904				
182					
183	Type species. Promyalina hindi Kittl, 1904; Permian, Changhsingian; near Sarajevo, Bosnia-				
184	Herzegovina.				
185	Diagnosis. Outline mytiliform and less triangular than Myalina; dorsal margin slightly				
186	convex, moderately thickened; narrow beak present. Left valve inflated below the umbo.				
187	Inequivalve with the right valve being less inflated and less curved.				
188					
189	Species <i>Promyalina schamarae</i> (Bittner, 1899)				
190	(Fig. 2C-G)				
191	1899 Myalina schamarae Bittner. p. 19; Plate 4; Figs 20-25.				
192	2009 Promyalina schamarae; Kumagae & Nakazawa. p. 157; Fig 144, 4-5.				
193	Material. Three specimens from LD-04 (NHMUK PE PEI 5484; NHMUK PI MB 1202;				
194	NHMUK PI MB 1205) and seven specimens from LD-05 (NHMUK PI MB 1188; NHMUK				

195	<u>PI MB 1203-1204; NHMUK PE PEI 5475-5476; NHMUK PE PEI 5500; NHMUK PE PEI </u>
196	<u>5503</u>).
197	Description. Shell is small, mytiliform, inequilateral, prosocline, higher than long, and
198	moderately inflated. Umbo is small, terminal and prosogyrate. Posterior dorsal margin is
199	straight. Posterior margin is slightly convex, forming a rounded posteroventral margin.
200	Anterior margin is long, nearly straight or weakly acute and partly depressed near the umbo.
201	Remarks. The internal characters of the shells cannot be observed because all the specimens
202	are articulated. Externally, the shells resemble $Promyalina\ schamarae$ from the Griesbachian
203	Lazurnaya Bay Formation, Russia (Shigeta et al. 2009), and are therefore assigned to this
204	species. These specimens differ from P. groenlandica (Newell 1955), as the beak does not
205	project beyond the dorsal margin, and from P. spathi in having a more convex anterior
206	margin.
207	Mode of life. Surficial, stationary, attached, suspension feeder (Stanley 1972).
208	
209	Order Solemyoida Dall, 1889
210	Family Nucinellidae Vokes, 1956
211	
212	Diagnosis. Shell nuculoid, obliquely oval, higher than long, monomyarian with anterior
213	abductor muscle scar only. Hinge with subumbonal taxodont teeth and single elongate lateral
214	tooth on the anterior dorsal margin. Ligament mostly opisthodetic, wholly external or in a
215	sunken resilifer.
216	Remarks. Although living nucinellids are sometimes classified in the family Manzanellidae,
217	which extends back into the Permian (e.g. Coan & Valentich-Scott 2012), Oliver & Taylor
218	(2012) argued that the Nucinellidae and Manzanellidae should be separated on morphological
219	grounds and their conclusions are followed here. Manzanella, the type genus of
220	Manzanellidae, is dimyarian and subcircular in outline, with its taxodont teeth lying posterior
221	to beak. In contrast, <i>Nucinella</i> and <i>Huxleyia</i> (i.e. the Nucinellidae) are both monomyarian and
222	elliptical, with their teeth positioned anterior to the beak (Oliver & Taylor 2012).
223 224	Genus <i>Nucinella</i> Wood, 1851
225	Type species. <i>Pleurodon ovalis</i> Wood, 1840; Neogene, Pliocene; Suffolk, England.
226	Diagnosis. Nucinellids with external ligament.

227	Remarks. Of the two genera described for the Nucinellidae, Huxleyia has a mostly internal
228	ligament set in a sunken resilifer, whereas Nucinella has an external ligament. The specimens
229	described herein have an opisthodetic or amphidetic ligament and so are assigned to
230	Nucinella.
231 232 233 234	Species <i>Nucinella taylori</i> sp. nov. (Fig. 3)
235	Diagnosis. A small <i>Nucinella</i> having a nuculoid shape, smooth shell except for growth lines
236	with three subumbonal and two anterior pointed blade like teeth, with no triangular flat area
237	below the dentition; opisthodetic ligament.
238	Holotype. Disarticulated Left valve, <u>NHMUK PI MB 1206</u> , LD-04; length =1.6mm, height =
239	1.6mm.
240	Paratypes. Disarticulated Right valve, <u>NHMUK PI MB 1209</u> , LD-04; length = 1.6mm,
241	height = 1.6mm; Disarticulated Right valve, <u>NHMUK PI MB 1210</u> , LD-05; length = 2.1mm,
242	height = 2.3mm.
243	Other Material. Two specimens from LD-04 (NHMUK PI MB 1207-1208) and three
244	specimens from LD-05 (NHMUK PE PEI 5500; NHMUK PE PEI 5506; NHMUK PE PEI
245	5517). Fifty-three prodissoconch valves from LD-04 (NHMUK PI MB 1211-1217; NHMUK
246	PE PEI 5481) and thirty-six prodissoconch valves from LD-05 (NHMUK PI MB 1218).
247	Derivation of name. Named after Dr. John Taylor (Natural History Museum) in recognition
248	of his work on chemosymbiotic bivalve molluscs.
249	Description. Shell is small, thin, inequilateral, equivalve, with a nuculoid, suboval outline.
250	Posterior dorsal margin is slightly incurved; anterior margin almost straight. Opisthogyrate,
251	beaks close to posterior margin, umbo prominent. Smooth shell, except for concentric,
252	irregularly spaced growth lines. Monomyarian: posterior adductor muscle scar absent;
253	anterior adductor large, oval. Opisthodetic ligament, prominent, external does not invade the
254	hinge plate. Hinge made by five pointed blade-like teeth as a single arched series, with the

Prodissoconch valves: outline nuculoid and suboval. Posterior margin is distinct, long and slightly incurved. Inequilateral, with beaks close to posterior margin, and sculpture consists of irregularly spaced growth laminae. Ventral valve margin has a narrow flat

anterior subumbonal teeth being smaller and wider than the posterior ones. A single, long

lateral tooth. Left valve with a secondary ridge creating a shallow socket.

platform. Five anterior and eleven posterior hinge teeth. Amphidetic ligament, lying between the beak and anterior hinge plate.

Remarks. These specimens are most similar to the extant <u>nucinellid</u> *Nucinella serrei*, in <u>their</u> small size, number of posterior and anterior hinge teeth and opisthodetic ligament. These specimens, however, lack a flat triangular area below the teeth dentition and a small circular pit at the end of the lateral tooth, which supports their separation.

The majority of extant *Nucinella* range from intertidal to 500m deep (la Perna 2005), but some species have been described from water depths exceeding 3000m (Oliver & Taylor 2012). A large fossil *Nucinella* species has been described from a Late Cretaceous cold-seep deposit (Amano et al. 2007), showing that this genus may inhabit a wide range of sulphiderich environmental settings. Bacterial symbiosis with sulphur-oxidizing bacteria is confirmed for *N. owenensis* and has been inferred for all species of the Nucinellidae (Oliver & Taylor 2012). *Nucinella taylori* sp. nov. supplants *N. birkelundi* from the Late Jurassic (Clausen & Wignall 1990) and *Nucinella*? sp. from the Late Triassic (Nützel & Kaim 2014) as the oldest known species of *Nucinella*, and extends the range of the genus to the basal Triassic *H. parvus* Conodont Zone.

The prodissoconch valves are very similar to adult specimens of *Nucinella taylori* sp. nov. except that they have more hinge teeth, which appears to reflect their pre-metamorphosis stage of development. The position of the ligament in the prodissoconch valves also differs to adult specimens of *N. taylori* sp. nov. in being amphidetic rather than opisthodetic, but this character is known to change after metamorphosis (Bernard, 1898). The prodissoconch valves most resemble *N. taylori* sp. nov. rather than *N. nakremi* sp. nov., but may represent larval stages for either or both species.

Mode of Life. Shallow infaunal, fully motile, slow, chemosymbiotic (Oliver & Taylor 2012).

Species <u>Nucinella</u> nakremi sp. nov.

286 (Fig. 4)

- Diagnosis: A small *Nucinella* having a nuculoid shape, smooth shell except for growth lines.
- 288 Prosogyrate beak, one to three subumbonal teeth. Ligament amphidetic and does not invade
- the hinge plate.
- **Holotype.** Disarticulated left valve, NHMUK PI MB 1219, LD-04; length = 1.1mm, height =
- 291 1.1mm.

support their separation.

- Paratype. Disarticulated left valve, <u>NHMUK PI MB 1220</u>, LD-04; length = 0.9mm, height = 0.8mm (transposed hinge).
- Other Material. Two specimens from LD-04 (NHMUK PI MB 1221-1222).
- Derivation of name. Named after Dr. Hans Arne Nakrem in recognition of his work on
- 296 Permian and Triassic fossils from Svalbard.
- **Description.** Shell small, nuculoid and ovate. Posterior dorsal margin distinct, slightly
- incurved; posterior margin rounded. Inequilateral, prosogyrate, with beaks close to anterior
- 299 margin. Umbo prominent. Smooth, thin shell with very weak growth lines. Monomyarian:
- posterior adductor muscle scar absent; anterior adductor large, oval. One to two subumbonal,
- 301 pointed blade-like teeth plus one anterior tooth. Ligament amphidetic, prominent, external
- does not invade the hinge plate. One left valve (LV) specimen with a long lateral tooth, and

Remarks. These specimens differ from other described nucinellid species in having fewer

- another LV specimen with a weak secondary ridge creating a shallow socket.
 - hinge teeth, a more elliptical shape and a prosogyrate beak. Such differences may occur during the ontogeny of *Nucinella* (e.g. Bernard, 1898) and so are not sufficient for assignment to a separate genus. These specimens are, however, considered to represent a separate species rather than an intermediate ontogenetic stage between the protoconch and adult stage of *Nucinella taylori*. During ontogeny, the shape, size and position of nucinellid subombonal teeth also vary: in earlier stages of development they are more rounded and later they develop a chevron-blade shape with the older teeth making space below the beak for thinner, newer ones (Bernard, 1898; La Perna, 2004). In contrast, the subumbonal teeth of specimens assigned to *N. nakremi* and *N. taylori* have comparable shapes, and so indicate a similar stage of development. Furthermore, because the position of the ligament in nucinellids is fixed after metamorphosis (Bernard, 1898; La Perna 2004), and the only subsequent ontogenetic change is an increase in ligament size with age, the differences in ligament position between *N. nakremi* and *N. taylori* cannot be ontogenetic. If the specimens assigned to *N. nakremi* were included as an intermediate ontogenetic stage of *N. taylori*, the ontogenetic pattern would not match any known nucinellid (cf. Bernard, 1898; La Perna,
 - A specimen of <u>N.</u> *nakremi* sp. nov. includes an example of a transposed hinge on a left valve (Fig. 4D). Instead of the normal left valve arrangement of three hinge teeth and a

2004), and, therefore, the differences in the subumbonal teeth and the position of the ligament

324	lateral secondary ridge creating a secondary socket, this specimen has two hinge teeth and a
325	lateral tooth (i.e. the normal right valve arrangement). An alternative interpretation is that this
326	specimen represents an earlier ontogenetic stage, but this is rejected because even though the
327	specimen is slightly smaller, its lateral tooth is more prominent than in a typical left valve
328	arrangement and both the subumbonal and lateral teeth appear to fit with the expected
329	corresponding valve arrangement. In addition, it has been reported that in the early stages of
330	nucinellid development the second lateral tooth is very small and closer to the subumbonal
331	teeth than later in ontogeny (Bernard, 1898), which is not the case with this specimen.
332	Transposed hinges have been reported in a number of bivalve families, but this is the first
333	reported occurrence in a species of Nucinellidae.
334	Mode of Life. Shallow infaunal, fully motile, slow, chemosymbiotic (Oliver & Taylor, 2012).
335 336	Order Nuculanoida Carter et al., 2000
337	Family Neilonellidae Schileyko, 1989
338	Diagnosis. Hinge plate with taxodont teeth in two series, sometimes separated by narrow,
339	plain area, without resilifer; ligament predominantly external, opisthodetic to amphidetic,
340	weak.
341	Remarks. The Neilonellidae are very similar to the Nuculanidae, but the lack of a resilifer in
342	mature adults supports their separation (Coan & Valentich-Scott 2012). The new specimens
343	described herein differ from the Malletiidae in lacking conspicuous gapes, and from the
344	Tindariidae in having a short gap in the dentition below the beaks (Di Geronimo & La Perna
345	1997). The family is currently known from the Jurassic to present in all oceans, especially in
346	deep water and soft substrates (Coan & Valentich-Scott 2012).
347	
348	Genus Austrotindaria Fleming, 1948
349	Type species. Austrotindaria wrighti Fleming 1948; Holocene, Quaternary; Southland, New
350	Zealand.
351	Diagnosis. Delicate almost smooth shell, with weak concentric sculpture. A short, edentulous
352	gap between the posterior and anterior hinge plates. No pallial sinus and without rostrum.
353	Remarks. The Neilonellidae comprises three valid genera: Neilonella, Austrotindaria and
354	Pseudoneilonella (La Perna 2007). The convex shape of the posterior margin and the

presence of an opisthodetic ligament in our specimens indicate better agreement with Austrotindaria than Neilonella (Di Geronimo & La Perna 1997; La Perna 2007). Austrotindaria differs from Neilonella and Pseudoneilonella by having a delicate, almost smooth, shell rather than a sturdy, sculptured shell, and by having a small elongate pit below the posterior and anterior teeth (La Perna 2007). This genus is currently known from the Miocene to Recent (La Perna 2007; Coan & Valentich-Scott 2012) and the specimens identified in his study extend the range of the genus to the beginning of the Triassic.

- Species Austrotindaria antiqua sp. nov.
- 364 (Fig. 5)
- 365 <u>1899 Anodontophora (Myacites) fassaensis; Bittner. p. 22; Plate. III, Figs 28-33.</u>
- 366 <u>1908 Anoplophoria (Myacites) fassaensis; von Wittenburg p. 33; Fig. 15.</u>
- 367 <u>1926 Anodontophora fassaensis</u>; Matsushita Plate 8; Fig. 11
- 368 1963 Unionites fassaensis; Ciriacks. p. 81; Plate 16; Fig. 13.
- 369 1985 Unionites fassaensis; var. brevis Neri & Posenato p. 94; Plate 2; Fig 8.
- 2009 Unionites fassaensis; Kumagae & Nakazawa p. 166; Fig. 145, 5-9.
- 371 2014 Unionites? Fassaensis; Pan et al. p. 151; Fig. 5, G-H.
- 372 2015 Unionites fassaensis; Hofmann et al. p. 8; Fig. 4, K.
- 2015 cf. Unionites fassaensis; Foster et al. p. 381, Fig. 4, L.
- Diagnosis. Shell small, smooth except for concentric growth lines, subtrigonal, inequilateral,
- prosogyrate; taxodont hinge dentition with more than 50% fewer anterior than posterior ones,
- obtusely chevron-shaped teeth, hinge plate interrupted below the beak by an edentulous gap,
- a small rounded triangular pit below the edentulous gap; ligament opisthodetic.
- Holotype. Disarticulated right valve, NHMUK PI MB 1240, LD-04; length = 4.0mm, height
- 379 = 3.0mm.
- **Paratype.** Disarticulated left valve, NHMUK PI MB 1241, LD-04; length = 4.6mm, height =
- 381 3.4mm.
- 382 Other Material. Eight specimens from LD-04 (NHMUK PI MB 1194; NHMUK PE PEI
- 383 5479; NHMUK PE PEI 5486-5488; NHMUK PE PEI 5513; NHMUK PE PEI 5520) and

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- eighteen specimens from LD-05 (NHMUK PI MB 1191; NHMUK PI MB 1196-1197;
- 385 NHMUK PI MB 1199; NHMUK PI MB 1242; NHMUK PE PEI 5475-5476; NHMUK PE
- 386 <u>PEI 5483</u>; <u>NHMUK PE PEI 5494</u>; <u>NHMUK PE PEI 5497</u>; <u>NHMUK PE PEI 5499-5501</u>;
- 387 NHMUK PE PEI 5504; NHMUK PE PEI 5506; NHMUK PE PEI 5515). Two prodissoconch
- 388 valves from LD-04 (NHMUK PI MB 1243-1244).
- Derivation of name. Latin, antiqua (ancient) referring to this species being the oldest known
- of the genus.
- **Description.** Equivalve, inequilateral shell with a subtrigonal outline and a beak positioned
- approximately 30% along the length of the dorsal margin from the posterior; H/L ratio 0.6-1.
- Conspicuously tumid. Umbo prosogyrate, prominent, moderately broad, rounded and projects
- 394 above the hinge margin. Slightly rounded and gently sloping anterodorsal margin.
- Posterodorsal margin almost straight to slightly rounded, gently sloping, with a slight angled
- junction with the posterior margin. Ventral margin deeply rounded. Escutcheon short,
- 397 relatively broad, elliptical; lunule narrow. Shell smooth with fine concentric growth lines;
- entire inner margin smooth. Small ligament, external, opisthodetic, with a well-defined
- margin and well-rounded triangular pit seated beneath the edentulous gap. Hinge plate with
- 400 taxodont teeth in two series, sometimes separated by a narrow, plain area, without resilifer,
- that is narrow below the beak, broadening towards the anterior and posterior ends. Robust
- 402 teeth, moderately long and blunt, with more than 50% fewer anterior than posterior ones.
- 403 Smooth ventral margin. No pallial line, sinus or muscle scars are present. As the size of the
- shell increases the edentulous gap becomes proportionally smaller, more central and moves
- 405 externally.
- **Remarks.** The external morphology of these specimens is identical to most Early Triassic
- specimens previously described as "Unionites" fassaensis, which is one of the most
- 408 widespread bivalve species from the Lower Triassic and is also a problematic dustbin taxon
- 409 that includes a range of different morphologies. These specimens differ from the original
- description and figures of "Myacites" fassaensis Wissmann (1841) in having a less elongated
- 411 posterior margin and a more prosogyrate beak. Due to poor preservation, however, little is
- known about the internal morphology of Early Triassic specimens assigned to *U. fassaensis*,
- which has created some uncertainty (Hautmann et al. 2013). Internally, these new specimens
- lack the following characters that Geyer et al. (2005) determined were diagnostic of
- 415 Unionites: a nymph extending nearly half of the posterior margin; an impressed adductor
- muscle scar; an overlap of the anterior hinge; and a deeply impressed lunule and posterior

keel. Thus, these specimens cannot be assigned to *Unionites*. The presence of taxodont hinge dentition, an opisthodetic ligament and lack of ornamentation is, however, characteristic of the genus *Austrotindaria*.

The more prosogyrate beak means that these specimens have a similar external morphology to Middle Triassic *Unionites* specimens (e.g. Geyer et al. 2005), but apart from possessing a faint posterior keel they lack the diagnostic criteria of *Unionites*. Furthermore, the *Unionites* specimens described by Geyer et al. (2005) and from this study (see below) all have a strong posterior keel. Thus, these specimens are not assigned to *Unionites*. These specimens, and all previous specimens assigned to *Unionites fassaensis* that possess a prosogyrate beak and lack a well-defined posterior keel, are therefore assigned to *Austrotindaria antiqua* sp. nov.

Mode of Life. Shallow infaunal, fully motile, slow, miner (Stanley 1968).

Species Austrotindaria svalbardensis sp. nov.

- 431 (Fig. 6)
- 432 <u>1864 Anoplophora fassaensis</u>; Alberti. p. 137; Plate III; Fig. 8.
- 433 1930 *Nucula* sp. juv. ind; Spath. p. 53; Plate XII; Fig 12.
- 434 2013b *Unionites fassaensis*; Hofmann et al. p. 887; Fig. 8, 17-18.
- 435 2015 Unionites cf. fassaensis; Hautmann et al. p. 890; Fig. 10, H.
- **Diagnosis.** Shell small, smooth except for concentric growth lines, subtrigonal, inequilateral,
- orthogyrate; taxodont hinge dentition with more posterior than anterior teeth, obtusely
- chevron-shaped teeth, hinge plate interrupted below the beak by an edentulous gap, a small
- rounded triangular pit below the edentulous gap; ligament opisthodetic.
- **Holotype.** Disarticulated right valve, NHMUK PI MB 1223, LD-04; length = 1.8mm, height
- 441 = 1.3mm.
- **Paratype.** Disarticulated left valve, NHMUK PI MB 1231, LD-04; length = 1.5mm, height =
- 443 1.0mm.
- Other Material. Twenty-one specimens from LD-04 (NHMUK PI MB 1190; NHMUK PI
- 445 MB 1192-1193; NHMUK PI MB 1224-1230; NHMUK PI MB 1232-1233; NHMUK PI MB
- 446 1235-1236; NHMUK PE PEI 5478; NHMUK PE PEI 5482; NHMUK PE PEI 5489;

- 447 <u>NHMUK PE PEI 5505; NHMUK PE PEI 5509-5510; NHMUK PE PEI 5512; NHMUK PE</u>
- PEI 5520) and twelve specimens from LD-05 (NHMUK PI MB 1189; NHMUK PI MB 1197;
- 449 NHMUK PI MB 1200; NHMUK PE PEI 5476; NHMUK PE PEI 5491-5492; NHMUK PE
- 450 <u>PEI 5496; NHMUK PE PEI 5498; NHMUK PE PEI 5503-5504; NHMUK PE PEI 5517</u>).
- Twenty-four prodissoconch valves from LD-04 (NHMUK PI MB 1234; NHMUK PI MB
- 452 <u>1237; NHMUK PI MB 1238; NHMUK PE PEI 5520</u>) and five prodissoconch valves from
- 453 LD-05 (<u>NHMUK PI MB 1239</u>).
- **Derivation of name.** Named after the Svalbard archipelago.
- Description. Shell outline subtrigonal, equivalve, inequilateral with beak positioned
- approximately 60% of the distance along the dorsal margin from the posterior; H/L ratio 0.6-
- 1; conspicuously tumid. Orthogyrate umbo, prominent, moderately broad, rounded and
- 458 projected above the hinge margin. Anterodorsal margin slightly rounded and gently sloping.
- The posterodorsal margin is almost straight to slightly rounded and gently sloping. Ventral
- 460 margin is smooth and deeply rounded. Escutcheon is short, relatively broad, elliptical; lunule
- arrow. No sculpture except for fine concentric growth lines; entire inner margin is smooth.
- 462 Prodissoconch is smooth, broadly subovate, with H/L ratio of 0.6-0.9. Ligament small,
- 463 external with well-defined margin, opisthodetic, with a well-rounded triangular pit seated
- beneath the edentulous gap. Hinge plate with taxodont teeth in two series; separated by
- arrow, plain area, without resilifer, narrow below the beak, broadening towards the anterior
- and posterior ends. Teeth are robust, moderately long and blunt, with more posterior than
- anterior teeth, separated by an edentulous gap. As the size of the shell increases the
- 468 edentulous gap becomes proportionally smaller, more central and moves externally. No
- pallial line, sinus or muscle scars present.
- **Remarks.** These specimens differ from Austrotindaria antiqua sp. nov. in having a more
- equal number of anterior to posterior teeth and an orthogyrate beak. The direction of the beak
- also separates these specimens from unequivocal species of *Unionites*. In the specimens
- 473 described here and other *Austrotindaria* species the beak is orthogyrate to posteriorly
- opisthogyrate, whereas in *Unionites* it is prosogyrate. Based on this key character, most
- 475 specimens previously assigned to *Unionites fassaensis*, and similar Early Triassic specimens
- with an orthogyrate beak and no internal morphological detail preserved, are herein assigned
- 477 to Austrotindaria svalbardensis sp. nov.
- These specimens have a very similar shape, size and ornamentation to the type species
- of Austrotindaria (A. wrighti Flemming), however, they differ in possessing a small, rounded,

triangular pit below the edentulous gap and differ from other Austrotindaria species, e.g. A. mawheraensis, in lacking a weak posterior rostrum. For these reasons, they are therefore considered a separate species. Mode of Life. Shallow infaunal, fully motile, slow, miner (Stasek 1961; Stanley 1968). Species Austrotindaria? canalensis (Catullo, 1846) (Fig. 2B) *Tellina canalensis* Catullo p. 56; Plate 4; Fig. 4. Anodontophora (Myacites) canalensis; Bittner p.23; Plate III; Figs 34-38. 1908 Anoplophoria canalensis; von Wittenburg p.281; Plate 5; Fig. 6. *Unionites canalensis*; Ciriacks p. 81; Plate 16; Fig. 11-12. Unionites canalensis; Neri & Posenato, p. 94; Plate 2; Fig. 8. 2004 Unionites aff. canalensis; Kashiyama & Oji p. 214; Fig 8, F. Unionites canalensis; Kumagae & Nakazawa p. 166; Fig. 145, 1-4. 2015 Unionites canalensis; Hofmann et al. p 481. Fig. 4, J. 2015 cf. Unionites canalensis; Foster et al. p. 381; Fig. 4, K. Material. Three specimens from LD-05 (NHMUK PI MB 1198; NHMUK PE PEI 5493; NHMUK PE PEI 5500). **Description.** Outline sub-ovate to elongate, equivalve, inflated below the umbo; inequilateral with beak lying approximately 62% along the dorsal margin length from posterior; H/L ratio 0.4-0.7. Posterior margin elongated and almost straight, anterior margin narrowly rounded. Escutcheon and lunule indistinct. Umbo orthogyrate, prominent, moderately broad, rounded, projecting above the hinge margin. Ornamented externally with fine concentric growth lines; entire inner margin smooth. **Remarks.** The internal characters and hinge in these specimens were not observed and have not been reported for *Unionites canalensis*. Externally, the shell is virtually identical to those that are typically assigned to *U. canalensis* (e.g. Hofmann et al. 2015), with a medially placed umbo that is a diagnostic feature of *U. canalensis* (Catullo, 1846) and an orthogyrate beak. In contrast, all other species of *Unionites* have a more anteriorly located umbo, and the beak in Unionites is prosogyrate (Geyer et al. 2005). Thus, these Early Triassic specimens clearly do not belong to the genus Unionites. The external features of these specimens are most similar to species of Neilonellidae, e.g. Austrotindaria benthicola (Dell, 1956). Austrotindaria is the only genus of the Neilonellidae that is reported from the Early Triassic, and these specimens are, therefore, tentatively assigned to it. Without observation of the internal characters an unequivocal generic assignment cannot be made. The posterior margin is more elongate than in Austrotindaria svalbardensis sp. nov., and these specimens are therefore considered to represent a separate species.

Mode of Life. Shallow infaunal, fully motile, slow, miner (Stanley 1968).

Family Malletiidae Adams & Adams, 1858

(Fig. 7)

- **Diagnosis.** Subequilteral to inequilateral; usually with elongate, compressed posterior end; sculpture of commarginal striae or ribs; anterior and posterior gapes present; hinge plate weak, with fine taxodont teeth in two series, sometimes separated by plain area, without resilifer; ligament external, opisthodetic to amphidetic, weak; pallial sinus large.
- **Material.** Three larval shells from LD-04 (NHMUK PI MB 1248-1250).
- Description. Shell is equilateral, elliptical, H/L ratio of 0.7, and moderately inflated. Umbo is broad, orthogyrate, with beak positioned centrally. Shell smooth except for concentric growth lines. Hinge plate has three anterior and three posterior teeth in two series separated by a large plain area with a groove. Ligament is predominantly external, amphidetic and weak.
 - Remarks. The Malletiidae are very similar to the Nuculanidae, but the lack of a resilifer in mature adults supports their separation (Coan & Valentich-Scott 2012). These specimens do not belong to the Neilonellidae because they possess a small conspicuous gape between the valves. The Malletiidae is a long-ranging family known from the Ordovician to the Recent, and three genera (Malletia, Palaeoneilo and Taimyrodon) belonging to the family have previously been reported from the Lower Triassic (e.g. He et al. 2007; Wasmer et al. 2012). Externally, these specimens resemble the larval shells of Paleoneilo? fortistriata figured by Wasmer et al. (2012), but they have far fewer hinge teeth. There are also equal numbers of anterior to posterior hinge teeth in these specimens, whereas in P. fortistriata there are many more posterior than anterior teeth (Wasmer et al. 2012).
 - **Mode of Life.** Shallow infaunal, fully motile, slow, miner (Stanley 1968).

541	
542	Order Unionoida Stoliczka, 1871
543	Family Anthracosiidae Amalitsky, 1892
544	Genus <i>Unionites</i> Wissmann, 1841
545	Type species. Unionites muensteri Wissmann 1841; Carnian, Late Triassic; South Tyrol,
546	Italy.
547	Diagnosis. The generic diagnosis follows Geyer et al. (2005). Shell suboval to elongate-
548	elliptical or rectangular to trapeziform; equivalve, inequilateral. Umbo prosogyrate. Shell
549	surface smooth except for growth rugae, or with rather fine commarginal riblets and, rarely, a
550	faint radial ornamentation. Anterior hinge margin of right valve projects beyond the plane of
551	commissure, fitting into a recess in the opposite valve. Small subumbonal groove limits this
552	shell projection posteriorly. Posterior lateral tooth of right valve generally distinct, engaging
553	below posterior dorsal margin of left valve. Hinge of left valve with anterior platform that
554	bears depression for corresponding anterior hinge margin of right valve, followed posteriorly
555	by small, tuberculiform subumbonal tooth. Narrow groove extends parallel to margin at distal
556	part of anterior hinge margin and corresponds to anterior lateral tooth of right valve. Lower
557	part of posterior dorsal margin slightly projects beyond plane of commissure, fitting above
558	posterior lateral tooth of right valve. Ligament fixed on nymph, which extends about half the
559	length of the posterior dorsal margin.
560	Species Unionites aff. subrectus (Bittner, 1901)
561	(Fig. 8)
562	aff. 1901 Anodontophora subrecta Bittner, p. 100; Plate 7; Figs. 28-30.
563	aff. 2003 Unionites subrectus; Szente & Vörös, p. 131; Plate Biv-II; Figs. 25-26.
564	2009 Triaphorus aff. multiformis; Kumagae & Nakazawa p. 171; Fig. 145, 18-25.
565	Material. Seven specimens from LD-04 (NHMUK PI MB 1245-1247; NHMUK PE PEI
566	5480; NHMUK PE PEI 5485; NHMUK PE PEI 5507; NHMUK PE PEI 5519) and nine
567	specimens from LD-05 (NHMUK PI MB 1195; NHMUK PI MB 1197; NHMUK PI MB
568	1201; NHMUK PE PEI 5495; NHMUK PE PEI 5506; NHMUK PE PEI 5508; NHMUK PE
569	<u>PEI 5516</u>).

Description. Outline elongate-elliptical. Shell equivalved, inflated below the umbo, and inequilateral with beak positioned approximately 85% along the length of the dorsal margin from the posterior. Lower part of anterior dorsal margin projects slightly beyond plane of commissure. Posterodorsal margin is almost straight to slightly round and gently sloping. Lunule is long and narrow, with deeply impressed escutcheon. Umbo is prosogyrate and rises above the hinge margin. Ornamented with fine concentric growth lines. The entire inner margin is smooth. Prodissoconch is smooth and orbicular.

A small subumbonal groove limits the shell projection posteriorly. Short lateral tooth, left valve possesses a weak secondary ridge creating a shallow socket. Hinge of left valve has an anterior platform that bears a depression for the corresponding anterior hinge margin of the right valve, posterior to this is a small, tuberculiform subumbonal tooth fitting above posterior lateral tooth of right valve. Ligament is fixed to a nymph, which extends about half the length of the posterior dorsal margin. Isomyarian muscle scars with a deeply impressed anterior adductor muscle scar.

Remarks. Insufficient knowledge of the internal morphology of Early Triassic bivalves in general, and *Unionites* in particular, has created uncertainty regarding their systematic position. Based on their external morphology alone these specimens would be assigned to *Triaphorus* aff. *multiformis* (Kumagae & Nakazawa 2009), but they possess the following characters that are diagnostic of *Unionites*: the anterior hinge margin of the right valve overlaps that of the left; a nymph that extends nearly half the length of the posterior hinge margin; an impressed anterior adductor muscle scar; a deeply impressed lunule; and a long and narrow escutcheon. Thus, these specimens are assigned to *Unionites*.

These specimens differ from *U. brevis*, *U. fassaensis*, *U. letticus*, *U. griesbachi*, *U. borealis*, and *U. rhomboidalis* because their length greatly exceeds their height. They are most similar to *U. subrectus* described from the Middle Triassic of Hungary by Szente & Vörös (2003), except they have a more elliptical outline.

Mode of Life. Shallow infaunal, facultatively motile, unattached, suspension feeder (Hautmann et al. 2013).

Order Trigoniida Dall, 1889

Family **Myophoriidae** Bronn, 1849

Genus *Neoschizodus* Gieberl, 1855

- Type species. *Lyrodon laevigatum* (Goldfuss, 1837), Anisian, Middle Triassic, Rüdersdorf,
- 603 Germany.
- **Diagnosis.** The diagnosis follows Newell & Boyd (1975). Orthogyrous to moderately
- prosogyrous shell with slightly incurved beak. Posterior ridge angular to subangular in
- transverse profile. Myophorian hinge; teeth and sockets smooth or bearing transverse
- 607 striations.

- Remarks. These specimens have most of the required diagnostic criteria for assignment to
- Neoschizodus: an orthogyrate to moderately prosogyrate shell with an incurved beak and a
- posterior ridge angular to subangular in transverse. The nymph, teeth and umbonal platform
- were not present in these specimens so it is not known whether these specimens possess the
- myophorian hinge that is also diagnostic of this genus. These specimens do, however, have a
- similar external morphology to other Early and Middle Triassic specimens assigned to
- Neoschizodus (e.g. Kumagae & Nakazawa 2009) and are, therefore, referred to this genus.

Species Neoschizodus laevigatus Gieberl, 1855

(Fig. 9)

- Material. Five specimens from LD-04 (NHMUK PI MB 1251; NHMUK PE PEI 5489;
- NHMUK PE PEI 5519; NHMUK PE PEI 5522) and eight specimens from LD-05 (NHMUK)
- 620 PE PEI 5476-5477; NHMUK PE PEI 5499; NHMUK PE PEI 5501; NHMUK PE PEI 5518;
- 621 NHMUK PE PEI 5521). Fifty-nine prodissoconch valves from LD-04 (NHMUK PI MB 1252;
- 622 NHMUK PI MB 1254; NHMUK PI MB 1255-1260; NHMUK PE PEI 5480; NHMUK PE
- 623 PEI 5484; NHMUK PE PEI 5486; NHMUK PE PEI 5511; NHMUK PE PEI 5513; NHMUK
- 624 PE PEI 5520) and fourteen prodissoconch valves from LD-05 (NHMUK PI MB 1253;
- 625 NHMUK PE PEI 5514).
- **Description.** Shell is trigonally subovate, equivalve, inflated below the umbo, inequilateral,
- slightly higher than long. Umbo is small and orthogyrate, with an elevated, subangular to
- 628 rounded, umbonal ridge. Indistinct posterior ridge. Anterodorsal margin recurvate, passing to
- 629 widely arched ventral margin, posterodorsal margin straight, entire inner margin smooth.
- 630 Smooth sculpture except for faint concentric growth lines.
- The larval shells have a similar external morphology to the adult shells, with a short
- and subumbonal hinge and short, narrow nymphs running down the anterior and posterior
- margins. The prodissoconch is smooth with fine concentric growth lines.

Remarks. Neoschizodus laevigatus is a cosmopolitan Early and Middle Triassic species with
a high variability in its morphological characters. The shell shape and indistinct posterior
ridge means that the present specimens agree with the characters of N . $laevigatus$, and similar
specimens identified from the Early Triassic (e.g. Kumagae & Nakazawa 2009; Hautmann et
al. 2011).

One of the adult specimens preserves a prodissoconch (Fig. 9I) and its morphology is identical to the other prodissoconchs that were found as isolated specimens. The hinge plate of the prodissoconchs was not observed to be myophorian, instead the hinge dentition is typical of a taxodont, but it is not known how a myophorian hinge plate develops through ontogeny (Newell & Boyd 1975). Hautmann & Nützel (2005) suggest that, in bivalves, the presence of a small prodissoconch I and a relatively large prodissoconch II indicates a planktotrophic larval stage. Therefore, a planktotrophic larval stage is interpreted for these specimens.

Mode of Life. Shallow infaunal, facultatively motile, unattached, suspension feeder (Hautmann et al. 2013).

Class Gastropoda Cuvier, 1795

Order **Amphigastropoda** Simroth, 1906

Family **Euphemitidae** Knight, 1956

653 Genus *Warthia* Waagen, 1880

- Type species. Warthia brevisinunata Waagen, 1880; Middle Permian, Salt Range, Pakistan.
- Diagnosis. The generic diagnosis follows Waterhouse (1963). Deeply involute planispirally
- 656 coiled bellerophont gastropod, distinguished by absence of spiral ornament. Sinus is broad,
- 657 usually with a short slit. Selenizone obscured by inductural layer. Ornament of faint growth
- 658 lines and growth wrinkles. Umbilicus generally filled.

Species Warthia zakharovi Kaim, 2009

660 (Fig. 10)

- 661 2009 Warthia zakharovi Kaim, p. 141; Fig. 132-134.
- Material. One-hundred and thirty-eight specimens from LD-04 (NHMUK PI MG 1468-1473);
- 663 NHMUK PI MG 1495-1496; NHMUK PI MG 1502-1511; NHMUK PE PEI 5478-5482;

NHMUK PE PEI 5484-5490; NHMUK PE PEI 5505; NHMUK PE PEI 5507; NHMUK PE PEI 5509-5513; NHMUK PE PEI 5519-5520; NHMUK PE PEI 5522) and one-hundred and ninety-three specimens from LD-05 (NHMUK PI MG 1467; NHMUK PI MG 1474-1481; NHMUK PI MG 1483; NHMUK PI MG 1485-1494; NHMUK PI MG 1497-1499; NHMUK PI MG 1512; NHMUK PE PEI 5475-5477; NHMUK PE PEI 5483; NHMUK PE PEI 5491-5504; NHMUK PE PEI 5506; NHMUK PE PEI 5508; NHMUK PE PEI 5514-5518; NHMUK PE PEI 5521). **Diagnosis.** Warthia with wide shell in comparison to its thickness and weakly depressed selenizone. **Description.** Shell globular, almost as long as wide. Slit short and broad at base of U-shaped sinus. Weakly depressed selenizone. In well-preserved specimens growth lines can be observed. Aperture is arched around earlier whorls, and curved inwards by the selenizone. **Remarks.** The shell of *Warthia* is usually entirely involute and overgrows the umbilicus. In some of our smaller specimens, the umbilicus is still visible, but in larger specimens it has become overgrown during subsequent growth. Even though our specimens lack an inductural layer and have growth lines that have not previously been recorded in Early Triassic Warthia species, in Permian species of Warthia visible growth lines or wrinkles have been recorded (e.g. W. micromphala and W. stricta; Waterhouse 1963) and, therefore, we consider our specimens to belong to Warthia. The question of whether bellerophontids had a planktonic larval stage is unresolved (Nützel & Mapes 2001). The protoconch in bellerophontids may be very small, less than one whorl, and is succeeded by the teleoconch (Frýda 1999), and in our specimens the initial whorl is overgrown almost immediately. The small size (<0.05mm), bilateral symmetry and lack of ornamentation in the initial whorls of Warthia zakharovi recorded in this study (Fig. 10H), are comparable to the embryonic shells of Bellerophon from the late Silurian that Frýda (1999) interpreted as indicating planktotrophy. A planktotrophic larval stage is, therefore, <u>inferred</u> for Warthia zakharovi. **Mode of Life.** Surficial, fully motile, slow, surface deposit feeder (Linsley 1977). Order Vetigastropoda Salvini-Plawen, 1980

Subfamily **Eotomariinae** Wenz, 1938

Family **Eotomariidae** Wenz, 1938

696	Genus Glabrocingulum Thomas, 1940
697	Type species. Glabrocingulum beggi Thomas, 1940, Carboniferous, Scotland.
698	Diagnosis. Low- to moderately high-spired and turbiniform shell shape. The upper whorl
699	surface forming an angle of <45° with the selenizone located on the upper edge of whorl face.
700	Sutures sharply defined. Upper whorl face with both spiral and collabral ornament; most
701	strongly developed near the suture, weakest near the selenizone. Anomphalus to widely
702	phaneromphalus, with or without funicle.
703	Remarks. These specimens resemble the Permian genera Wannerispira, Ananias and
704	Glabrocingulum, and the Triassic genus Kamupena. They differ from Wannerispira by
705	possessing a selenizone in the upper third of the whorl, and only having two rather than three
706	strong spiral ribs; from Ananias by being low- rather than high-spired and having a less
707	conspicuous and thinner concave band below the selenizone; and from Kamupena by lacking
708	a strong umbilical callus plug. These specimens also differ from other neilsoniines by having
709	spiral ribs, no axial ornamentation and by being less elongated. Another genus with a
710	comparable whorl profile is Rhaphistomella, which has been considered as a synonym of
711	Glabrocingulum (Batten 1989; Erwin & Pan 1996), but it differs from these specimens by the
712	absence of a prominent medial concave band and a more strongly nodulose keel under the
713	suture. These specimens are therefore assigned to the genus Glabrocingulum.
714	Wannerispira is the only other unequivocal eotomariid genus to have been reported
715	from the Early Triassic (Kaim et al. 2010; Hautmann et al. 2015) and belongs to the
716	Subfamily Neilsoniinae. Since Glabrocingulum, in contrast, is dextral, low-rather than high-
717	spired, and with a moderately deep slit developing into a selenizone with rounded margins, it
718	belongs within the Subfamily Eotomariinae. These specimens represent the first Early
719	Triassic record of the Subfamily Eotomariinae and are the first Early Triassic record of the
720	genus Glabrocingulum, which is rarely recorded after the Permian period.
721	
722	Species Glabrocingulum parvum sp. nov.
723	(Fig. 11)
724	Diagnosis. Moderately low-spired shells with faint spiral ornament without sutural nodes.
725	Broad selenizone in relation to whorl height with sharp edges. Deeply concave selenizone.
726	Funicle absent or weakly formed.
727	Holotype. Dextral shell, NHMUK PI MG 1531; height = 2.5mm, width = 3.0mm.

- 728 Paratype. Dextral shell, NHMUK PI MG 1518; height = 2.0mm, width = 2.0mm.
- 729 Other Material. Ten specimens from LD-04 (NHMUK PI MG 1468; NHMUK PI MG 1471;
- 730 NHMUK PI MG 1495; NHMUK PI MG 1500-1501; NHMUK PI MG 1513-1515; NHMUK
- 731 PE PEI 5519-5520) and fourteen specimens from LD-05 (NHMUK PI MG 1476; NHMUK
- 732 <u>PI MG 1482; NHMUK PI MG 1484; NHMUK PI MG 1493; NHMUK PI MG 1530;</u>
- 733 NHMUK PE PEI 5493; NHMUK PE PEI 5502; NHMUK PE PEI 5514). Fourteen juvenile
- 734 shells from LD-04 (<u>NHMUK PI MG 1516-1517</u>; <u>NHMUK PI MG 1519-1529</u>; <u>NHMUK PE</u>
- 735 <u>PEI 5486</u>).
- **Derivation of name.** Latin, parvum (small), referring to its small size.
- **Description.** Shell <u>is dextral</u>, turbiniform, low-spired, <u>with simple sutures</u>. The upper whorl
- surface is slightly concave and gently sloping, and bears the selenizone between sharply
- 739 <u>protruding edges, with the lower of the edges on the</u> shell periphery. Selenizone is concave
- and moderately deep. Narrow, concave band immediately below the lower rib. Whorl profile
- below the lower rib is gently convex. Basal angulation is relatively sharply defined, but
- convex; base with a rounded circum-umbilical shoulder; small umbilical chink. Aperture is a
- rounded trapezoid; inner lip is reflexed; peristome interrupted by a slit in the outer lip. Shell
- ornamented with closely, irregular spaced fine spiral lirae. Growth lines visible with small
- knobs at the intersection of spiral ribs, otherwise no axial ornamentation observed. On the top
- of the keel, near the suture, the growth lines form small nodules.
- Protoconch: openly coiled; first two whorls smooth; third whorl possesses ~15 evenly
- spaced, rounded, spiral threads; peristome uninterrupted (Fig. 11F).
- **Remarks.** These specimens <u>resemble</u> Glabrocingulum texanum Batten (1989) with the
- selenizone being located in the upper third of the whorl and lacking axial ornamentation.
- 752 They differ from G. texanum in having a broader selenizone in relation to whorl height; a
- more concave selenizone with sharper edges; a weakly developed funicle or none at all; and
- in being moderately low-spired. The uncoiling that has been described for some G. texanum
- specimens from the Permian of the southwestern US (Batten 1989) was not observed. These
- specimens are also considerably smaller (Max size H=3.6mm, W=4.6mm) than the type
- material of G. texanum (Max size H=8.7mm, W=9.9mm), which may be a consequence of
- 758 environmental stress in the immediate aftermath of extinction event and an expression of the
- 759 <u>Lilliput effect in this genus (cf. Twitchett 2007)</u>. Due to their excellent preservation, these

specimens reveal the morphology of the larval stages, which show a similar ontogenetic development as other species of *Glabrocingulum* (e.g. Pan and Shen, 2008).

The <u>Vetigastropoda</u> have a diverse range of living habits including being described in association with wood-fall communities (Kiel et al. 2008). The specimens in this study do occur in association with wood, however, no direct relationship was observed. Eotomariidae <u>recorded</u> from Zechstein <u>reefs</u> are described as motile algal grazers that were probably confined to a hard substrate (Hollingworth & Pettigrew 1988). <u>The lack of evidence</u> of a hard substrate in <u>their depositional environment</u>, however, suggests that these specimens probably had a similar life habit to deep-sea vetigastropods that typically consume sediment (Hickman 1988).

Mode of Life. Surficial, fully motile, slow, surface deposit feeder.

Order Ptenoglossa Gray, 1853

Family **Pseudozygopleuridae** Knight, 1930

774 (Fig. 1<mark>2</mark>)

Material. Three larval shells from LD-04 (NHMUK PI MG 1532-1534).

Description. Teleoconch not observed. Protoconch is conical, elongate and composed of five whorls. The initial whorl is smooth and has a diameter of 0.08mm. Collabral ornamentation is initiated on the second whorl, continues to the base of the protoconch, and consists of narrow costellae that intersect at or just below the mid-whorl. Costellae are sigmoidal: on the upper part of the whorls they are slightly curved and oriented 40° to the shell axis, and on the lower part of the whorls they are oriented 200° to the shell axis. Growth lines are visible as faint collabral ribs between, and perpendicular to, the costellae. Aperture is circular, with a small columellar fold. Four whorls are present in these specimens.

Remarks. Hoare & Sturgeon (1978) showed that protoconchs of species of Pseudozygopleuridae are very similar, but can be readily differentiated from those of the Zygopleuridae. Diagnostic characters of pseudozygopleurid protoconchs that are present in these specimens include an elongate, conical shape with 3-5 whorls; smooth initial whorl with collabral ornamentation from the second whorl; and narrow opisthocline to sigmoidal transverse costellae that are equally spaced, extend up and below from the suture and curve uniformly to the mid-whorl (Hoare & Sturgeon 1978). In contrast, protoconchs of the

Zygopleuridae have smooth whorls with fine riblets at the sutures (Nützel & Mapes 2001; Kaim 2004) or straight ribs in the Ampezzopleurinae (Nützel 1998; 2005). Protoconchs of the Ladinulidae are similar to those of the Pseudozygopleuridae but can be differentiated due to their vertical costellae (Bandel 2006). These specimens are, therefore, representatives of the Pseudozygopleuridae. No teleoconch is apparently present in these specimens as the sculpture of the prodissoconch is uninterrupted.

The Pseudozygopleuridae is primarily a Palaeozoic family. Apart from the Early Triassic specimens described herein, the only other Mesozoic pseudozygopleurid is *Plocezyga* from the Jurassic of Poland (Kaim 2004). Those specimens attributed to *Plocezyga* by Kaim (2004), however, lack a protoconch with the diagnostic sigmoidal pseudozygopleurid ornamentation described by Hoare & Sturgeon (1978; 1980). Instead they have a reticulate pattern, shouldered whorls and transverse costae on the protoconch, indicating that they belong to a different genus and family.

Nützel (1998) used the diameter, ornamentation and whorl number of the protoconch) to separate planktotrophic from non-planktotrophic pseudozygopleurid species. The specimens recorded in this study have small protoconch I diameters (0.12-0.14mm), and the same ornamentation and number of whorls as pseudozgopleurid larval shells interpreted as planktotrophic by Mapes & Nützel (2009), which suggest that they too had a planktotrophic larval lifestyle. A teleoconch was not developed in any of the specimens, which may indicate that the adults were unable to live in this environment.

Order Architectibranchia Haszprunar, 1985

Family **Tubiferidae** Cossmann, 1895

814 Genus *Sinuarbullina* Gründel, 1997

Type species. Sinuarbullina ansorgi Gründel, 1997; Jurassic, Bathonian, NE Germany, NW

816 Poland.

Diagnosis. The generic diagnosis follows Gründel and Nützel (2012). The shell is fusiform with a distinctly elevated spire. The teleoconch whorls have a subsutural ramp. The transition from the ramp to the outer whorl face is either rounded or angular, sometimes demarcated abapically by a spiral concavity. Whorls are smooth in most species, although faint spiral furrows occur on the base in a few of them. Growth lines are prosocyrt on the outer whorl face but strongly curving in an abapertural direction and opisthocyrt at the ramp. The aperture

is relatively low for the group, teardrop-shaped and lack columnellar fold. The protoconch is trans- to medioaxial.

Remarks. Pan et al. (2003) described *Jiangxispira* as a new genus using the following diagnostic features: 1) a small, high-spired, slender, fusiform shell; 2) teleoconch whorls with narrow shoulder; and 3) a heterostrophic, discoidal, transaxial protoconch. These characters are all within the diagnostic features of the genus *Sinuarbullina* Gründel (1997), and *Jiangxispira* is, therefore, a junior synonym of *Sinuarbullina*.

Species *Sinuarbullina yangouensis* (Pan et al., 2003) comb. nov. (Fig. 13)

2003 Jiangxispira yangouensis Pan et al. p. 44; Fig. 3, 1-7.

Material. Specimen lost by WJF after photography.

Description. The shell is high-spired, slender and fusiform. Teleoconch whorls have a subsutural ramp. The ramp from the outer whorl face is rounded with a rib on the shell periphery. Whorls are smooth, except for growth lines which are prosocyrt on the outer whorl face curving in an apertural direction and become opisthocyrt towards the ramp. The surface of the shell shows a colour spiral band around the subsutural ramp. The aperture is an elongated teardrop shape. Protoconch is heterostrophic, sinistral, nearly discoidal with lightly elevated spire 30° offset from the shell axis; protoconch has 1-2 round whorls.

Remarks. Seven species are included in *Sinuarbullina*, and *S. convexa* (='Cylindrobullina' convexa) is the only accepted species from the Lower Triassic (Gründel & Nützel 2012). These specimens are more slender than *S. convexa*, described from the Sinbad Limestone of the western US by Batten & Stokes (1986), and better resemble *Jiangxispira yangouensis* from the Induan Dayie Formation, China.

The shell morphology is similar to *Meekospira*, which has been interpreted as a slow-moving shell dragger (Hughes 1986), but could have also been a burrower (Hollingworth & Pettigrew 1988). Interpreting the feeding strategy of fossil gastropods is difficult because information on the organ system, including the ctendium, is not usually preserved. The ancestral condition in high-spired gastropods is presumably an algal grazer on hard substrates (Declerck 1995). Given the absence of hard substrates in this study, the specimens described herein were probably detritus feeders or possibly micro-carnivorous on sedentary prey as in many modern shelled opisthobranchs (e.g. Lobo da Cunha et al. 2009).

Mode of Life. Surficial, fully motile, slow, deposit feeder.

Discussion

Implications for Permian-Triassic extinction and diversification

The fully silicified, diverse, earliest Triassic assemblage described herein includes gastropod and bivalve taxa that have not previously been recognised from the Early Triassic. Some occurrences extend stratigraphic ranges back to the basal Triassic (Fig. 14), with implications for timing and rates of diversification, whereas others show unequivocally that some taxa that were previously thought to have become extinct during the late Permian mass extinction event actually survived.

Three of the four gastropod taxa represent families that originated in the Palaeozoic, of which two (*Warthia* and *Wannerispira*) have previously been recognised from the earliest Griesbachian and have been interpreted as 'dead clades walking' (*sensu* Jablonski 2002) by Kaim et al. (2010) and Kaim & Nützel (2011). The family Pseudozygopleuridae is recorded in the Triassic for the first time (Fig. 14), confirming that it too survived the late Permian extinction, and could also be interpreted in this way. As noted above, the Jurassic specimens assigned to the Pseudozygopleuridae by Kaim (2004) do not have the expected suite of diagnostic characters and it may be that they represent a different family. If so, then the specimens described herein from Svalbard would represent the youngest occurrence of the Pseudozygopleuridae.

The other gastropod described in this study, *Sinuarbullina yangouensis*, represents one of the oldest occurrences of the architectibranchs (*sensu* Gründel & Nützel 2012), a group that diversified in the late Triassic and Jurassic. Although architectibranchs have been described from the Carboniferous, the Carboniferous specimens do not have a heterostrophic protoconch, which means that they are not true architectibranchs (Gründel & Nützel 2012). Thus, while the oldest unequivocal architectibranchs occur in the first conodont zone of the Triassic (this study; Pan et al. 2003), as hypothesised for other benthic invertebrate groups, such as the articulate crinoids (Baumiller et al. 2010; Oji & Twitchett 2015), their origin may have been Palaeozoic.

Of the five bivalve genera identified in this study, *Promyalina*, *Unionites* and *Neoschizodus* have previously been recorded in the earliest Induan (e.g. Hautmann et al. 2015). The occurrences of the protobranchs *Nucinella* and *Austrotindaria* represent significant range extensions back to the Induan and are the oldest occurrences of the

Nucinellidae and Nuculanida respectively. Recent phylogenetic analyses show that the protobranchs rapidly diversified in the late Triassic and Jurassic (Bieler et al. 2014), but these new occurrences demonstrate that the Nucinellidae and Neilonellidae appeared ca. 50 million years before that, in the basal Triassic, which better supports the view of Sharma et al. (2013) that the late Permian mass extinction event triggered protobranch diversification. Future phylogenetic analyses will need to take into account our new findings in order to better calibrate molluscan evolutionary trees.

The excellent preservation of the fossil assemblage described in this study has been critical in recognising that <u>most</u> specimens previously assigned to *Unionites fassaensis* and *U. canalensis* (Palaeoheterodonta) from Lower Triassic rocks worldwide are likely to have been misidentified and better <u>resemble</u> the protobranch *Austrotindaria*. Other specimens from this study are, however, unequivocally identified as *Unionites*, confirming its presence in Svalbard during the earliest Triassic. The palaeoheterodonts *Unionites* and *Neoschizodus* represent two lineages that were part of a major Early Triassic diversification of the Palaeoheterodonta (Newell & Boyd 1975; Ros et al. 2011; Sharma et al. 2013; Bieler et al. 2014).

The re-assignment of most Early Triassic "Unionites" specimens to Austrotindaria has important implications for understanding how benthic marine ecosystems functioned in the wake of the late Permian mass extinction event, as pre-Spathian shelly benthic assemblages are typically described as being dominated by "Unionites" (Fraiser & Bottjer 2007; Hofmann et al. 2013b; 2014; 2015; Foster et al. 2015). Our taxonomic re-assignment has implications for the functional interpretation of those specimens previously described as "Unionites" because the motility and feeding of Unionites and Austrotindaria are significantly different, i.e. the former is a facultatively motile suspension-feeder and the latter is a motile deposit-feeder. Other non-protobranch taxa with a similar morphology, e.g. Tellina, also have a deposit-feeding mode of life, so this ecological reinterpretation will still be valid even if the tentative generic reassignment from "Unionites" canalensis to Austrotindaria? canalensis is subsequently revised. This new functional interpretation means that the palaeoecology of Early Triassic benthic shelly assemblages is in better agreement with the ichnofaunal record than previously thought (cf. Twitchett 2006), and that in many post-extinction, pre-Spathian benthic communities the dominant functional group was infaunal, motile, deposit-feeders.

The Middle and, especially, the Late Triassic are frequently described as being key intervals in the radiation of extant marine invertebrate groups from both fossil and molecular

evidence (e.g. Nakazawa & Runnegar 1973; Nützel 2005; Hautmann, 2007; Posenato 2008; Gründel & Nützel 2012; Rouse et al. 2013; Bieler et al. 2014; Hausmann & Nützel 2014). The poor quality of the Early Triassic fossil record is, however, widely recognised (Fraiser & Bottjer 2005) and it is demonstrably biased towards certain regions, palaeolatitudes and depositional settings (Foster & Twitchett 2014). As shown by other studies (e.g. Hautmann & Nützel 2005; Oji & Twitchett 2015), taxa may be overlooked because they are small, poorly preserved or inhabited particular depositional settings, and there is significant hidden diversity in the Early Triassic. The present study reinforces that view as the assemblage documented herein is the first fully silicified Early Triassic fauna, comes from a remote location, comprises mainly small-sized fossils, and demonstrates that Early Triassic diversity was higher than previously recognised.

The high-fidelity silicification, which has preserved shells that were originally aragonitic in exquisite detail, is a key factor. A significant post-Permian reduction in the number of silicified assemblages has been attributed to a decline and movement offshore of siliceous sponges (Schubert et al. 1997), driven by changes in climate, ocean circulation and productivity (Kidder & Erwin 2001). Although partially silicified assemblages are known from the Early Triassic (e.g. Twitchett et al. 2004), which Fraiser & Bottjer (2005) argued are adequate for palaeoecological analyses, the present study has demonstrated the necessity of early, complete and high-fidelity silicification for detailed taxonomic and ecological analysis. Thus, it appears that there is significant hidden biodiversity in the Early Triassic and that the diversification of many extant marine groups probably began earlier than is presently recognised, with implications for both the timing and rate of evolution. The most diverse silicified assemblages of the earliest Induan are known from offshore settings (Twitchett et al. 2004; this study), consistent with Schubert et al.'s (1997) hypothesis that siliceous sponges moved offshore after the late Permian extinction event, and so a search strategy focusing on those depositional settings is likely to yield critical new information, especially if regions that are currently under-represented are prioritised.

The importance of planktotrophy

The exceptionally preserved prodissoconchs and protoconchs in this study provide valuable insights into the early ontogeny of these taxa. Valentine & Jablonski (1983; 1986) suggested that during the late Permian extinction event there was selection against benthic invertebrates with planktotrophic larval stages. More recently, however, planktotrophic larval

shell development has been inferred for many Early Triassic gastropods (Nützel & Erwin 2002; Pan et al. 2003; Nützel & Schulbert 2005) and this selectivity has been questioned (Nützel 2014). All gastropod taxa recorded in the present study (i.e., *Warthia, Glabrocingulum, Sinuarbullina* and the Pseudozygopleuridae) are inferred to have had planktotrophic larvae (Nützel & Mapes 2001, this study), and in addition, the bivalves *Nucinella, Austrotindaria, Unionites* and *Neoschizodus* all possess a small prodissoconch I and a relatively large prodissoconch II which imply planktotrophic larval development (cf. Hautmann & Nützel 2005). Thus, we infer that the possession of planktotrophic larvae was a particular advantage for benthic molluscs inhabiting shelf settings of the Boreal Ocean in the earliest Triassic. This may simply be a consequence of sampling this particular depositional setting, as benthic taxa with plankotrophic larvae were common in similar mid-outer shelf settings during the Palaeozoic (e.g. Nützel & Mapes 2001; Frýda 2001; Bandel et al. 2002). Alternatively, it may be a consequence of the extinction event and associated environmental changes.

One possible advantage of a planktotrophic larval development is protection from benthic predation (Nützel & Frýda 2003). Although predators are rarely identified in Early Triassic benthic assemblages (e.g. Schubert & Bottjer 1995; Twitchett et al. 2004; Hautmann et al. 2011; Hofmann et al. 2013a; 2013b; 2014; 2015; Foster and Twitchett 2014; Foster et al. 2015), it is possible that some nektobenthic conodonts, ammonoids or fish may have fulfilled this role or that the predators were largely non-mineralised. Fish with durophagous dentitions, such as *Bobasatrania*, are locally common in the lowest Triassic of the Boreal Ocean (e.g. East Greenland; Stensiö 1932), and trace fossil evidence of vagile crustaceans has been recorded from the Induan of northwest Canada (Beatty et al. 2008; Zonneveld et al. 2010) and central Spitsbergen (WJF & RJT pers. obs.).

An alternative advantage is that planktotrophic larvae enable more effective dispersal and a wide geographic distribution. The cosmopolitan opportunists that thrived in the wake of the late Permian mass extinction (e.g. *Lingularia* and *Claraia*) all have an inferred planktotrophic larval stage (Hammond & Poiner 1984; Yang et al. 2001). Possession of planktotrophic larvae would have been particularly useful for Early Triassic benthic invertebrates given the spatial and temporal fluctuations in benthic oxygen concentrations that have been inferred for the earliest Triassic seafloor (Wignall & Twitchett 1996; Thomas et al. 2004; Nabbefeld et al. 2010). Effective dispersal and wide distribution would have afforded greater protection from extinction and enabled rapid colonisation of vacated seafloor once harsh environmental conditions had ameliorated. The late Permian mass extinction did

not select against animals with planktotrophic larval development (Posenato et al. 2014; Nützel 2014), and possession of that character may have been a key factor in the survival and radiation of certain groups.

The occurrence of well-preserved planktotrophic larval shells in this newly described benthic assemblage has additional palaeoenvironmental implications. First, it implies that primary production in surface waters was adequate enough to sustain planktic foodwebs. Second, the exquisite preservation of the earliest formed larval stages and the complete absence of any signs of shell dissolution or repair (cf. Garilli et al. 2015) demonstrates that neither the surface waters nor benthic habitat in this particular region were affected by ocean acidification at this time, and the pH was not low enough to have curtailed biomineralization.

Conclusions

The fossil assemblage from the Lusitaniadalen section, Svalbard, is the first fully silicified fauna to have been described from the Early Triassic and provides new critical systematic data. The fauna includes four new bivalve species: *Nucinella taylori*, *N. nakremi*, *Austrotindaria antiqua* and *A. svalbardensis* and a gastropod species: *Glabrocingulum parvum*. The silicified fauna from Svalbard demonstrates that the aftermath of the late Permian mass extinction was a key interval for the diversification of the Architectibranchs, Protobranchia and Palaeoheredonta. The assignment of Early Triassic specimens previously identified as *Unionites* to *Austrotindaria* has important palaeoecological implications demonstrating that infaunal deposit-feeders dominated benthic assemblages prior to the Spathian. The gastropod and bivalve taxa recorded in the present study are inferred to have had planktotrophic larvae, and it is likely that this was an important adaptation for bivalves and gastropods in mid-outer shelf settings of the Boreal Ocean during the earliest Triassic.

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- **Figure 1.** Locality and stratigraphy of the studied section in Lusitaniadalen. **A,** locality map
- of the Lusitaniadalen section; B, Palaeogeographic position; C, Stratigraphic column
- indicating the position of the sampled concretionary levels (LD-04 and LD-05) and
- disappearance of bioturbation associated with marine ecosystem collapse and the Late
- Permian mass extinction (LPE; modified from Nabbefeld et al., 2010). KSF = Kapp Starostin
- Formation, ii = ichnofabric index. Palaeogeography after Blakey (2012).

- 1380 Figure 2. A, Orbiculoidea winsnesi Gobbet (1963), NHMUK PI BE 3238, dorsal view,
- dorsal valve; B, Austrotindaria? canalensis (Catullo, 1846), NHMUK PI MB 1198, external
- view, left valve; C-G, *Promyalina schamarae* (Bittner, 1899), external view; C-D, NHMUK
- 1383 PI MB 1204; C, right valve; D, left valve. E-F, NHMUK PI MB 1203; E, right valve; F, left
- valve. G, NHMUK PI MB 1202, right valve. Scale bar: A-B, G = 2mm, C-F = 1mm.

- **Figure 3.** *Nucinella taylori* sp. nov. **A,** Holotype, Length 1.6mm, height 1.5mm,
- 1387 NHMUK PI MB 1206, external view, left valve; **B,** Paratype, Length 1.6mm, height –
- 1388 1.6mm, NHMUK PI MB 1209, external view, right valve; C, Paratype, Length 2.1mm,
- height 2.3mm, NHMUK PI MB 1210, internal view, right valve; **D,** Holotype, NHMUK PI
- 1390 MB 1206, internal view, left valve; E-F, Paratype, NHMUK PI MB 1209, internal view, right
- valve; G-J, Prodissoconchs of N. taylori: G, NHMUK PI MB 1213, external view, left valve;
- **H, NHMUK PI MB 1211**, external view, right valve; **I, NHMUK PI MB 1214**, internal view,
- left valve; **J, NHMUK PI MB 1211**, internal view, right valve. Scale bar: A-F = 500µm; G-J
- $1394 = 100 \, \mu \text{m}$

- **Figure 4.** *Nucinella nakremi* sp. nov. **A,** Holotype, Length 1.1mm, height 1.1mm,
- 1397 NHMUK PI MB 1219, external view, left valve; B, Paratype, Length 0.9mm, height –
- 0.8mm, NHMUK PI MB 1220, external view, left valve; C, Holotype, NHMUK PI MB 1219,
- interior view, left valve; **D**, Paratype, NHMUK PI MB 1220, interior view (transposed
- hinge), left valve. Scale bar = 500μ m.

- **Figure 5.** Austrotindaria antiqua sp. nov. **A,** Holotype, Length 4.0mm, height 3.0mm,
- 1403 NHMUK PI MB 1240, external view, right valve; B, Paratype, Length 4.6mm, height –
- 3.4mm, NHMUK PI MB 1241, external view, right valve; C, Holotype, NHMUK PI MB
- 1405 1240, internal view, right valve; **D, NHMUK PI MB 1191**, external view, right valve; **E,**
- 1406 NHMUK PI MB 1199, external view, left valve; F, NHMUK PI MB 1196, external view,

- right valve; G, NHMUK PI MB 1242, hinge plate of left valve; H, NHMUK PI MB 1194,
- dorsal view, right valve; I, Articulated specimen, NHMUK PI MB 1243; J, Larval shell,
- 1409 NHMUK PI MB 1244, left valve. Scale bar: A-C, $G = 500 \mu m$; D-F = 1mm; I-J = 100 μm .
- Figure 6. Austrotindaria svalbardensis sp. nov. A, Paratype, Length 1.5mm, height –
- 1.0mm, NHMUK PI MB 1231, external view, left valve; **B**, Holotype, Length 1.8mm,
- height 1.3mm, NHMUK PI MB 1223, external view, right valve; C, NHMUK PI MB 1190,
- 1414 <u>left valve, external view; **D**, larval shell, NHMUK PI MB 1234, external view, right valve; **E**,</u>
- 1415 NHMUK PI MB 1225, internal view, left valve; F, NHMUK PI MB 1227, internal view, left
- valve; G, larval shell, NHMUK PI MB 1237, internal view, right valve; H, Paratype,
- posterior and anterior hinge plate junction, <u>NHMUK PI MB 1231</u>. Scale bar = $500\mu m$.
- 1419 Figure 7. Malletiidae (Adams & Adams, 1858) larval shell. A, NHMUK PI MB 1249,
- external view; **B**, NHMUK PI MB 1249, internal view; **C**, NHMUK PI MB 1248, internal
- view. Scale bar = $100\mu m$.
- **Figure 8.** Unionites aff. subrectus (Bittner, 1901). A, NHMUK PI MB 1246, external view,
- right valve; **B, NHMUK PI MB 1245**, external view, left valve; **C, NHMUK PI MB 1246**,
- internal view, right valve; **D, NHMUK PI MB 1245**, internal view, left valve; **E, NHMUK PI**
- 1426 MB 1195, external view, right valve; F, hinge margin, NHMUK PI MB 1245, left valve; G,
- hinge margin, NHMUK PI MB 1246, right valve. Scale bar = 1mm.
- **Figure 9.** Neoschizodus laevigatus Gieberl (1855). A, NHMUK PI MB 1251, external view,
- left valve; **B-G**, **J**, larval shells; **B**, NHMUK PI MB 1256, external view, right valve; **C**,
- 1431 NHMUK PI MB 1259, external view, left valve; **D,** NHMUK PI MB 1257, external view,
- right valve; E, NHMUK PI MB 1260, internal view, left valve; F, NHMUK PI MB 1258,
- internal view, left valve; G, NHMUK PI MB 1255, internal view, right valve; H-I, beak of
- adult specimen, NHMUK PI MB 1251; J, hinge plate of a larval shell, NHMUK PI MB 1260.
- Scale bar: A = 2mm; B-G, $I = 200\mu m$; H = 1mm; $J = 50\mu m$.
- Figure 10. Warthia zakharovi Kaim (2009). A, NHMUK PI MG 1502, apertural view; B-C,
- 1438 NHMUK PI MG 1505; B, apertural view; C, lateral view; D-E, NHMUK PI MG 1509; D,
- lateral view; **E**, apertural view; **F**, NHMUK PI MG 1508, apical view; **G-H**, NHMUK PI MG
- 1440 <u>1504</u>; **G,** apical view; **H,** protoconch. A-B, Scale bar = 500 μ m, except H.

1441	
1442	Figure 11. <u>Glabrocingulum parvum sp. nov.</u> A-B, , apertural view; A, <u>Holotype, Height –</u>
1443	2.5mm, Width – 3.0mm, NHMUK PI MG 1531; B, Paratype, Height – 2.0mm, Width –
1444	2.0mm, NHMUK PI MG 1518; C-E, NHMUK PI MG 1531; C, lateral view; D, side view of
1445	initial whorls; E, apical view; F, apertural view of a larval shell, NHMUK PI MG 1520. Scale
1446	bar = 1mm, except F.
1447	
1448	Figure 12. Pseudozygopleuridae Knight (1930), larval shell. A-B, NHMUK PI MG 1532; A,
1449	apertural view; B , lateral view; C , NHMUK PI MG 1533, apical view; D , side view of initial
1450	whorl, NHMUK PI MG 1532. Scale bar = $100\mu m$.
1451	
1452	Figure 13. Sinuarbullina yangouensis (Pan et al., 2003). A, lateral view; B-C, apertural view;
1453	D , view of the protoconch. Scale bar = 1mm, except D. Note: Specimen lost by WJF after
1454	photography.
1455	
1456	Figure 14: Ranges of benthic invertebrate families identified from the base of the
1457	Vikinghøgda Formation. Grey bars = range extensions based on this study; Black bars =
1458	previously known unequivocal ranges of families; Dashed line = ghost ranges. Chang. =
1459	Changhsingian. Radiometric ages after: Lehrmann et al., 2006; Galfetti et al., 2007; Shen et
1460	al,. 2011. Late Permian mass extinction event is indicated by a vertical dark grey bar.

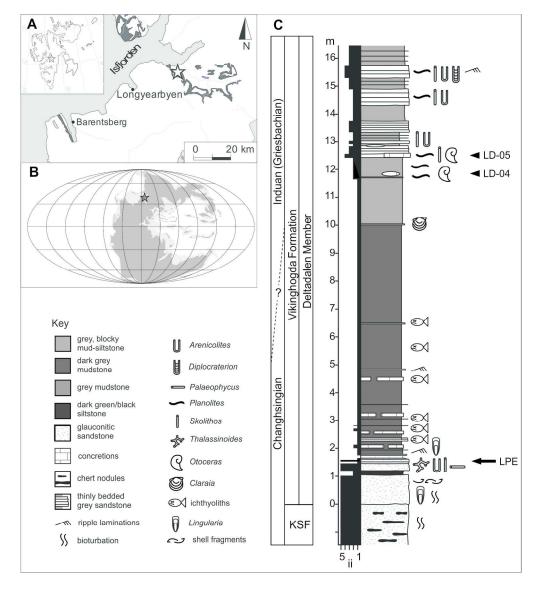


Figure 1. Locality and stratigraphy of the studied section in Lusitaniadalen. (A) locality map of the Lusitaniadalen section; (B) Palaeogeographic position and (C) Stratigraphic column indicating the position of the sampled concretionary levels (LD-04 and LD-05) and disappearance of bioturbation associated with marine ecosystem collapse and the Late Permian mass extinction (LPE; modified from Nabbefeld et al., 2010). KSF = Kapp Starostin Formation. ii = ichnofabric index. Palaeogeography after Blakey (2012).

179x200mm (300 x 300 DPI)

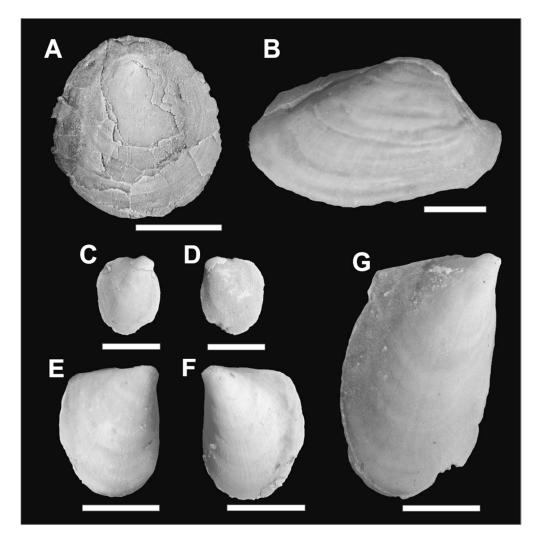


Figure 2. A, Orbiculoidea winsnesi Gobbet (1963), NHMUK PI BE 3238, dorsal view, dorsal valve; B, Austrotindaria? canalensis (Catullo, 1846), NHMUK PI MB 1198, external view, left valve; C-G, Promyalina schamarae (Bittner, 1899), external view; C-D, NHMUK PI MB 1204; C, right valve; D, left valve. E-F, NHMUK PI MB 1203; E, right valve; F, left valve. G, NHMUK PI MB 1202, right valve. Scale bar: A-B, G = 2mm, C-F = 1mm.

85x86mm (300 x 300 DPI)

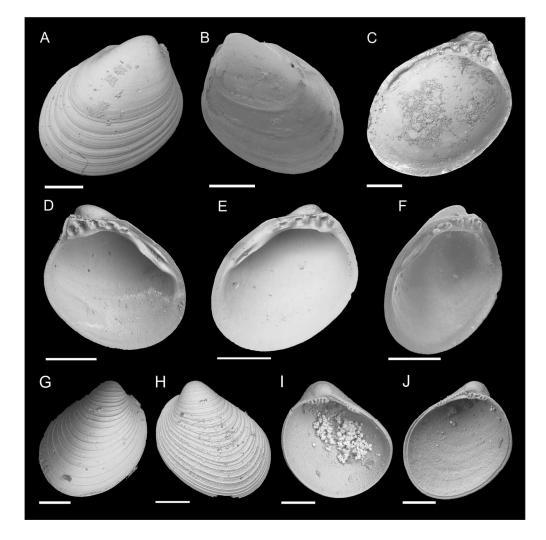


Figure 3. Nucinella taylori sp. nov. A, Holotype, Length – 1.6mm, height – 1.5mm, NHMUK PI MB 1206, external view, left valve; B, Paratype, Length – 1.6mm, height – 1.6mm, NHMUK PI MB 1209, external view, right valve; C, Paratype, Length – 2.1mm, height – 2.3mm, NHMUK PI MB 1210, internal view, right valve; D, Holotype, NHMUK PI MB 1206, internal view, left valve; E-F, Paratype, NHMUK PI MB 1209, internal view, right valve; G-J, Prodissoconchs of N. taylori: G, NHMUK PI MB 1213, external view, left valve; H, NHMUK PI MB 1211, external view, right valve; I, NHMUK PI MB 1214, internal view, left valve; J, NHMUK PI MB 1211, internal view, right valve. Scale bar: A-F = 500μm; G-J = 100 μm.

170x171mm (300 x 300 DPI)

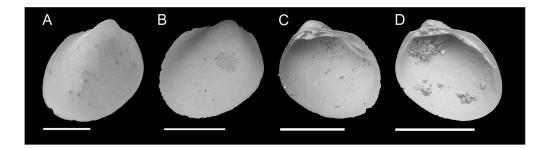
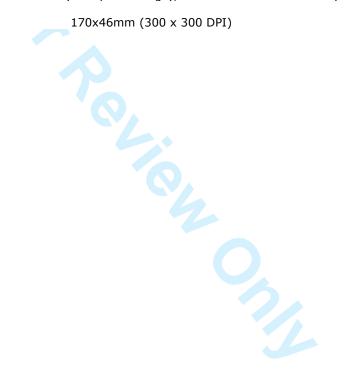


Figure 4. Nucinella nakremi sp. nov. A, Holotype, Length – 1.1mm, height – 1.1mm, NHMUK PI MB 1219, external view, left valve; B, , Paratype, Length – 0.9mm, height – 0.8mm, NHMUK PI MB 1220, external view, left valve; C, Holotype, NHMUK PI MB 1219, interior view, left valve; D, Paratype, NHMUK PI MB 1220, , interior view (transposed hinge), left valve. Scale bar = 500µm.



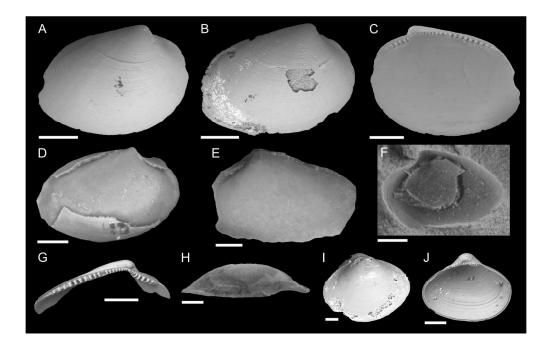


Figure 5. Austrotindaria antiqua sp. nov. A, Holotype, Length – 4.0mm, height – 3.0mm, NHMUK PI MB 1240, external view, right valve; B, Paratype, Length – 4.6mm, height – 3.4mm, NHMUK PI MB 1241, external view, right valve; C, Holotype, NHMUK PI MB 1240, internal view, right valve; D, NHMUK PI MB 1191, external view, right valve; E, NHMUK PI MB 1199, external view, left valve; F, NHMUK PI MB 1196, external view, right valve; G, NHMUK PI MB 1242, hinge plate of left valve; H, NHMUK PI MB 1194, dorsal view, right valve; I, Articulated specimen, NHMUK PI MB 1243; J, Larval shell, NHMUK PI MB 1244, left valve. Scale bar: A-C, G = 500μm; D-F = 1mm; I-J = 100μm.

170x108mm (300 x 300 DPI)

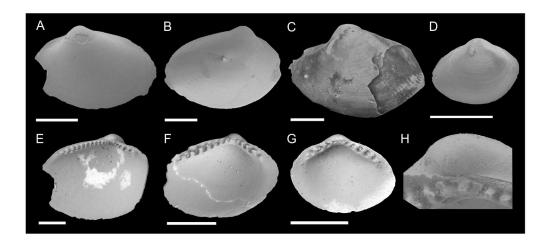


Figure 6. Austrotindaria svalbardensis sp. nov. A, Paratype, Length – 1.5mm, height – 1.0mm, NHMUK PI MB 1231, external view, left valve; B, Holotype, Length – 1.8mm, height – 1.3mm, NHMUK PI MB 1223, external view, right valve; C, NHMUK PI MB 1190, left valve, external view; D, larval shell, NHMUK PI MB 1234, external view, right valve; E, NHMUK PI MB 1225, internal view, left valve; F, NHMUK PI MB 1227, internal view, left valve; G, larval shell, NHMUK PI MB 1237, internal view, right valve; H, Paratype, posterior and anterior hinge plate junction, NHMUK PI MB 1231. Scale bar = 500μm.

170x75mm (300 x 300 DPI)

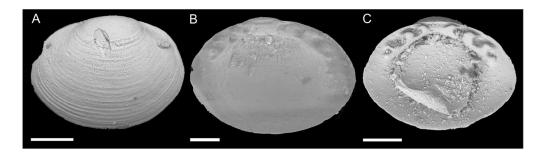


Figure 7. Malletiidae (Adams & Adams, 1858) larval shell. A, NHMUK PI MB 1249, external view; B, NHMUK PI MB 1249, internal view; C, NHMUK PI MB 1248, internal view. Scale bar = 100μm.

170x47mm (300 x 300 DPI)

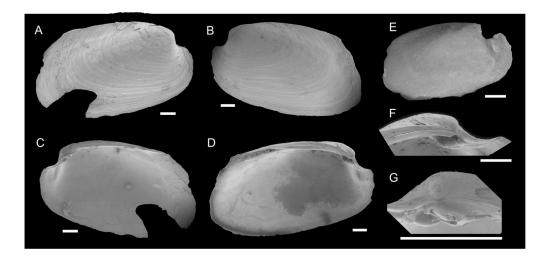


Figure 8. Unionites aff. subrectus (Bittner, 1901). A, NHMUK PI MB 1246, external view, right valve; B, NHMUK PI MB 1245, external view, left valve; C, NHMUK PI MB 1246, internal view, right valve; D, NHMUK PI MB 1245, internal view, left valve; E, NHMUK PI MB 1195, external view, right valve; F, hinge margin, NHMUK PI MB 1245, left valve; G, hinge margin, NHMUK PI MB 1246, right valve. Scale bar = 1mm.

85x40mm (600 x 600 DPI)

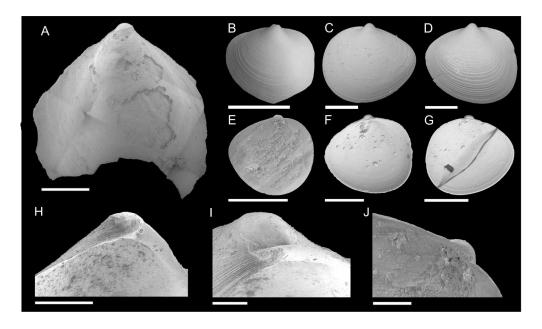


Figure 9. Neoschizodus laevigatus Gieberl (1855). A, NHMUK PI MB 1251, external view, left valve; B-G, J, larval shells; B, NHMUK PI MB 1256, external view, right valve; C, NHMUK PI MB 1259, external view, left valve; D, NHMUK PI MB 1257, external view, right valve; E, NHMUK PI MB 1260, internal view, left valve; F, NHMUK PI MB 1258, internal view, left valve; G, NHMUK PI MB 1255, internal view, right valve; H-I, beak of adult specimen, NHMUK PI MB 1251; J, hinge plate of a larval shell, NHMUK PI MB 1260. Scale bar: A = 2mm; B-G, I = $200\mu\text{m}$; H = 1mm; J = $50\mu\text{m}$.

170x100mm (300 x 300 DPI)

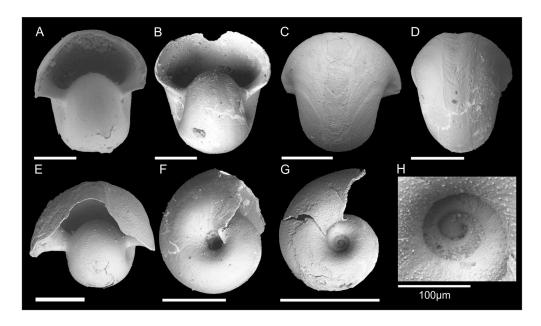


Figure 10. Warthia zakharovi Kaim (2009). A, NHMUK PI MG 1502, apertural view; B-C, NHMUK PI MG 1505; B, apertural view; C, lateral view; D-E, NHMUK PI MG 1509; D, lateral view; E, apertural view; F, NHMUK PI MG 1508, apical view; G-H, NHMUK PI MG 1504; G, apical view; H, protoconch. A-B, Scale bar = $500 \mu m$, except H.

170x100mm (300 x 300 DPI)

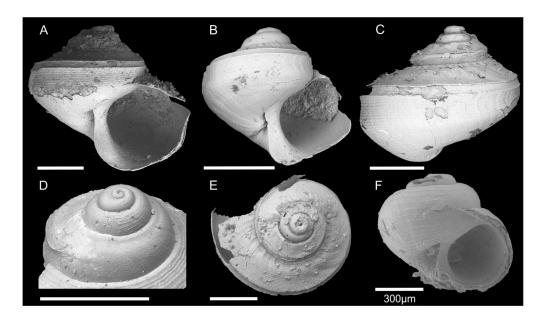


Figure 11. Glabrocingulum parvum sp. nov. A-B, , apertural view; A, Holotype, Height – 2.5mm, Width – 3.0mm, NHMUK PI MG 1531; B, Paratype, Height – 2.0mm, Width – 2.0mm, NHMUK PI MG 1518; C-E, NHMUK PI MG 1531; C, lateral view; D, side view of initial whorls; E, apical view; F, apertural view of a larval shell, NHMUK PI MG 1520. Scale bar = 1mm, except F.

170x98mm (300 x 300 DPI)

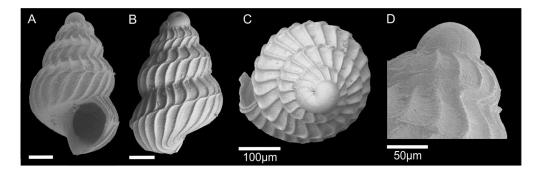


Figure 12. Pseudozygopleuridae Knight (1930), larval shell. A-B, NHMUK PI MG 1532; A, apertural view; B, lateral view; C, NHMUK PI MG 1533, apical view; D, side view of initial whorl, NHMUK PI MG 1532. Scale bar = 100µm.



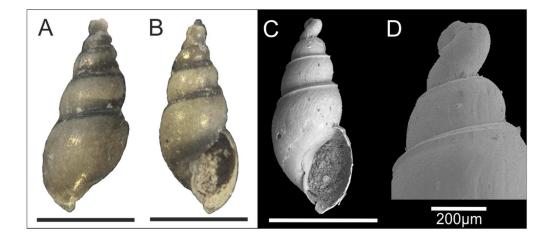


Figure 13. Sinuarbullina yangouensis (Pan et al., 2003). A, lateral view; B-C, apertural view; D, view of the protoconch. Scale bar = 1mm, except D. Note: Specimen lost by WJF after photography.

94x41mm (300 x 300 DPI)

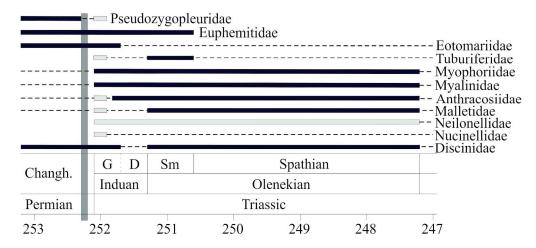


Figure 14: Ranges of benthic invertebrate families identified from the base of the Vikinghøgda Formation. Grey bars = range extensions based on this study; Black bars = previously known unequivocal ranges of families; Dashed line = ghost ranges. Chang. = Changhsingian. Radiometric ages after: Lehrmann et al., 2006; Galfetti et al., 2007; Shen et al,. 2011. Late Permian mass extinction event is indicated by a vertical dark grey bar.

199x88mm (300 x 300 DPI)