

Climate change and growing megacities: Hazards and vulnerability

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ABSTRACT

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The forms and frequency of geophysical and atmospheric hazards associated with extreme environmental conditions are evolving as global and regional climates change. Compounding effects due to the global growth of megacities modulate the impacts of such hazards. Rising carbon emissions, largely produced by increasing energy demand in urban areas throughout the world, is considered to be the main cause of long-term changes to the global climate. The correspondingly greater impacts of geophysical and atmospheric hazards within urban regions, where the majority of the world's inhabitants reside, raises concerns regarding vulnerability and response to varied pressures within the context of a changing climate. These hazards and their changing variability depend on geography, type of urbanization, and socio-economic factors. This paper is a review of current and anticipated geophysical and climatic trends associated with extreme weather events and natural hazards, their succeeding implications for urban areas, and the compounding pressures added by continued environmental modification due to urban expansion. We review how urban design, technological development, and societal behaviour can either ameliorate or worsen climate-induced hazards in urban areas. Pressures ranging from excessive rainfall leading to urban flooding, to positive and negative urban temperature extremes - which can simultaneously amplify or diminish local hazards such as locally and distantly produced high air pollution - require attention in order to understand, model, and predict changes in hazards in urban areas.

Keywords: Urban, hazards, vulnerability, sustainability, modelling, regional climate change.

53 **1. Introduction and Overview**

54 From the earliest times, socio-economic factors and the ease of withstanding natural hazards
55 in larger groups have led people to conglomerate. Throughout the world from the smallest
56 communities to the largest cities, these gatherings have formed self-replicating patterns of
57 organisation (Hunt 2005; Batty, 2008). However, as the Second United Nations Habitat
58 Conference in 1996 recognised (UN, 1996), cities, “especially in developing countries”, can
59 also be responsible for degradation of regional environments with harmful impacts for people
60 and ecosystems, including deterioration of health and safety, increased air pollution, inadequate
61 sewage and water management distribution systems, and modification of patterns and
62 intensities of storms and floods (Grimm et al., 2008; Hunt, 2009a; Hondula et al., 2014). Risks
63 from disease and pandemics coupled with increased exposure owing to population increase and
64 climate change also have implications for future vulnerability of urban areas (Hunt et al., 2016).

65
66 While cities serve as important agents that provide economic (e.g., employment), social (e.g.,
67 education), and a host of biophysical benefits (e.g., access to clean water and sanitation), their
68 increasing size also places undue strain on infrastructure, increases energy demand, and has
69 lead to ecological degradation (Grimm et al., 2008). As this paper further explains, the
70 increasing size and population of cities generally lead to worsening of environmental hazards.
71 In addition, these factors extend the distances around cities where hazards can be exacerbated.
72 This is why cities are also becoming more vulnerable to hazards produced by other megacities
73 located upwind and by upwind environmental dangers, such as smoke from burning and
74 pollutants from shipping (e.g. Zhang et al., 2011; Cheng and Chan, 2012; Lin et al., 2014; Li
75 et al., 2015). As cities and clusters of cities, or conurbations, expand, their energy use
76 (Madlener and Sunak, 2011) and pollution emissions increase (Akimoto, 2003). Cities can also
77 alter the adjoining rural environment, and the rivers and coasts that are so crucial to the
78 livelihoods of small communities and natural ecosystems (Ehrenfeld, 2000; Lee et al., 2006;
79 Shao et al., 2006; Aguilar, 2008; Georgescu et al., 2009; Salvati et al., 2012; Li et al., 2016a;
80 Yang et al., 2016a). In particular, although recent estimates indicate cities take up less than 1%
81 of the Earth’s landmass (Schneider et al., 2010), they are responsible for a disproportionate
82 amount of long-lived greenhouse gas emissions (UN, 2007; Satterthwaite, 2008).

83
84 This paper is firstly a review of current and projected extreme weather event trends and
85 associated natural hazards, as well as of most recent studies in the field. Secondly, we consider
86 the effects of these occurrences on urban hazards and their impact on urban environmental
87 vulnerability and sustainability. The conclusions highlight how studies of hazards and
88 vulnerability provide the basis for estimating impacts and policy development.

89 **2. Trends in climate related hazards**

90 Trends in climate related hazards will be discussed in two sub-sections focusing mainly on
91 their relation to regional climate extremes and urban growth.

92 **2.1. Extremes in regional climate and environment**

93 New observations and earth system models are showing how the climate has varied in the past
94 in different characteristic ways (IPCC, 2014). Weather patterns are determined by climate, but
95 also by orographic factors and other elements of the earth-climate system with their own
96 intrinsic variability (Schellnhuber et al., 2004). Over millennia, familiar atmospheric
97 circulations such as temperate westerly winds and sub-tropical trade winds have persisted, even
98 through ice ages (Houghton, 2015). However, although hemispheric oceanic circulations such
99 as the Gulf Stream have endured, there have been large fluctuations, affecting ocean
100 temperatures in sub-arctic regions (Broecker, 2010). Observational evidence reveals local
101 climatic effects associated with natural variations of atmospheric winds and ocean currents
102 over annual and decadal periods – for example, the El Nino Southern Oscillation (ENSO) –
103 and movements and variability of zones of forestation and desertification, such as the once in
104 a century southward movement of the Sahel in the 1970's.

105
106 Paleo-climate models, and bio-chemical measurements, now show that humans have also
107 played a role in influencing regional variability of climate, originally through agricultural and
108 forestry management, and more recently through development of built environments and
109 greenhouse gas emissions (Hunt, 2005; Lentz et al., 2014). For example, the magnitude of
110 current melting of Arctic summer ice and glaciers across the world is larger than what has
111 occurred naturally since the last ice ages about 10,000 years ago. The conclusion of IPCC
112 (2014) is that this trend is likely to be a result of human effects, although high levels of internal
113 variability can mask or interrupt the visibility of anthropogenically-induced changes (Swart et
114 al., 2015). A schematic diagram illustrating the diversity of natural effects and hazards
115 influenced by climate change is shown in Figure 1.

116
117 IPCC (2014) concludes with high confidence that globally averaged near-surface temperatures
118 will remain essentially constant for centuries even if anthropogenic emissions stopped
119 completely, unless there is a considerable net removal of carbon dioxide from the atmosphere.
120 Furthermore, continued decline of biological species associated with a changing climate is
121 likely to endure unless the current trend in climate change begins to reverse. Recent
122 measurements suggest that ice sheet temperatures may have already risen to the point where
123 polar and mountain ice sheets and glaciers are beginning to fracture and slide into the ocean at
124 a sufficient volume rate that they may continue to do so, even if the global average temperature
125 returns to 1850 values during the forthcoming 200-300 years. This would cause significant sea
126 level rise of several meters over the next millennium, resulting in catastrophic flooding and
127 associated impacts on global society (IPCC, 2014). The current scientific majority view is that
128 unless the future level of human influence on climate were to decline sometime during this
129 century through global action (therefore, halting the current rise of carbon emissions by or prior
130 to mid-century; see Stern, 2006), aspects of the climate system (e.g., biosphere) will begin to
131 change irreversibly (Lenton et al., 2008). This is the assumption of much current policy-making
132 and is the subject of this review. However, this notion is not shared by the entire scientific
133 community (Lawson, 2008). For example, while Solomon et al. (2009) argue that ice-cap loss,
134 hence sea level rise, is irreversible due to the longevity of atmospheric carbon dioxide

135 emissions and already risen ocean temperatures, Notz (2009) defends that there is no “tipping
136 point” for the loss of Arctic summer sea ice; therefore, sea ice is more likely to recover if
137 climate warming is stopped by reversing carbon emissions to pre-industrialization levels (if not
138 to prior), as supported also by Tietsche et al. (2011). Conversely, the irreversibility of the ice
139 sheet loss covering Greenland and the West Antarctic cannot be ruled out (Notz, 2009).

140
141 Recent analysis of climate prediction models and observational analysis indicate that, while
142 short-term trends (i.e. decadal scale or less) may not necessarily reveal long-term trends, the
143 effects of increasing carbon emissions have played an important role in global warming,
144 generally exceeding 0.1°C/decade, since the middle of the previous century (Tollefson, 2016).
145 Of key relevance for urban areas are effects on the variability of climate, including their impacts
146 on the global environment and society. Atypical events include ‘extreme’ events, defined
147 broadly as events that differ from average weather and climate, and/or that may persist over
148 longer periods. Other atypical events are the significant changes in trends, including those
149 owing to occurrence of extreme events (IPCC, 2012; Hov et al., 2013). There are significant
150 similarities between the main features of climate variability and other complex systems. An
151 important characteristic is that as fluctuations increase in frequency and magnitude, the non-
152 linear interactions between the various components of the processes under investigation
153 become more significant. In addition, physically based reductionist models become more
154 reliable than purely statistical extrapolations based on past events (Hunt et al., 1996; Hunt et
155 al., 2012).

156
157 An example of short-term, high magnitude events are strong convective updrafts and
158 downdrafts, resulting from higher surface temperature, deeper troposphere and cooler
159 stratosphere, which leads to higher rainfall intensity (now reaching 200 mm/hour and double
160 its value 10 years ago in south-east Asia; see Hong Kong Observatory, 2016; Wong et al., 2011;
161 Lee et al., 2010). Consequently, increased frequency of lightning has been observed (Hunt et
162 al., 2010; ten Hoeve et al., 2012) with a wider global distribution extending to Arctic regions,
163 as demonstrated by regular monitoring from satellites and globally from ground networks. This
164 trend is leading to enhanced fire risk, forest degradation and destruction, more rapid run-off
165 and consequential drought in some mountainous areas. These trends in convective storm events
166 are associated with, especially in southeastern USA, alteration in the changing nature of
167 tornadoes with increased maximum wind speeds and greater widening of the affected regions
168 (Hunt and Hangan, 2013; Elsner et al., 2015). However, observational analysis for other
169 portions of the USA indicates a decreasing or near-zero trend in tornado temporal variability
170 since 1950 (Guo et al., 2016), highlighting the importance of regionally based impact and
171 vulnerability analysis.

172
173 Equally important types of extremes include periods of very warm or cold weather, rainy or
174 snowy, or very windy, which occur more frequently and persist over considerably longer
175 periods than observed historically (WMO, 2013; IPCC, 2014; Heim, 2015; Matthews et al.,
176 2016). Model simulations combined with physical arguments and exploratory analysis of
177 recent global weather anomalies have concluded that such ‘blocking’ events will occur more

178 often and last longer (Cassou and Guilyardi, 2007; Li et al., 2012). Sometimes, these events
179 occur in a region simultaneously as climatic anomalies in other regions (Cheung et al., 2012).
180 A recent notable illustration was the extreme 2010 flooding in Pakistan, which was associated
181 with blocking in western Russia and persistent high temperatures in Moscow. Importantly, not
182 all global computer models have come to this conclusion (Pelly and Hoskins, 2003).

183
184 Global climate modelling has progressively become more useful (i.e., since the 1970's) as
185 spatial resolution has improved, leading to better representation of the key regional variations
186 of planetary climate. The incorporation of urban canopy models within earth system models
187 enables improved understanding of the interacting components of the earth system jointly with
188 human activities (e.g., Li et al., 2016b&c). Deficiencies in the representation of the marked
189 trends and fluctuations in regional climate remain, such as mountainous regions (with reduced
190 snowmelt and extremes in flooding), polar regions associated with ocean-atmosphere-
191 cryosphere interactions, and in modelling the lower tropical atmosphere, including the
192 particular effects of urban areas at high spatial resolution (Shaffer et al., 2016).

193 ***2.2. Climate change and urban growth leading to increased hazards and vulnerability***

194 Although the great cities of the world were largely founded for furthering trade, they were also
195 designed, in part, to protect their citizens from natural and artificial hazards including those
196 associated with extreme weather conditions (Hunt 2009b; Hunt, 2013). However, urban
197 hazards may become greater, as human activities change and the overall size of the city
198 (denoted by diameter L in Figure 2) grows. In Asia and Africa an additional 16% increase in
199 the total urban population is expected for both continents by 2050 (UN, 2014; for more
200 information on increase in urban population see
201 <http://data.worldbank.org/indicator/SP.URB.TOTL.IN.ZS>). This is a significantly faster rate
202 than that at which global climatic parameters are changing, such as the annual average
203 temperature increase relative to pre-industrial values (doubling over about 70 years) (for
204 Chinese megacities see Chan and Yao, 2008). These changes are likely to lead to increased
205 energy use and emissions of pollution. In addition, recent research has shown that projected
206 impacts on near-surface temperature within urban areas are of the same order of magnitude as
207 effects due to large-scale climate change (Georgescu et al., 2013, 2014), underscoring the
208 significance of cities as instruments of adaptation and mitigation.

209
210 Even with no change in the form of or the type of activities within a city, as the built up area
211 (i.e., size) and energy use increase, differences in temperature and humidity between the urban
212 and rural areas grow significantly. Atmospheric flows are changed as a result of increased
213 convection and turbulence. In desert areas (such as central China) strong inversion layers and
214 dust lead to trapping of air pollution. In addition, the larger the city, the larger the total
215 emissions of air pollutants and the higher the concentration (approximately in proportion to L ;
216 Figure 2). As their size increases above the characteristic 'meso-scale' distance, i.e. Rossby
217 radius L_R (Hunt et al., 2004), the influence of the earth's rotation becomes significant at mid-
218 latitudes. The physico-chemical properties of the air and water, and many aspects of the
219 biosphere, depend on the relative sizes of green and built-up areas (Figure 2). Changes in these

220 properties also depend on the buildings and planning of cities, their infrastructure, and people's
221 social behaviour patterns and government structures (e.g., Eakin et al., 2017), which vary
222 between regions and countries, and among large and small cities or even rural areas. In addition,
223 most people spend about 80% of their time indoors, where environmental factors (e.g.,
224 temperature, air circulation and quality, and humidity) are generally designed for much of the
225 time to be markedly different from those outside (Sailor, 2014; Rupp et al., 2015). Indoor
226 environmental control affects energy use and the actual levels of temperature and air quality
227 experienced by occupants (Hunt and Li, 2014), but have consequential implications for outdoor
228 environments (e.g., Taleghani et al., 2013; Salamanca et al., 2014).

229

230 In 'low-rise' megacities, the average heights of buildings are approximately constant (e.g., in
231 Europe, Africa, and some USA cities such as Phoenix, AZ), although most now have one or
232 more central business districts, which have incorporated buildings of greater vertical
233 dimension. However, in 'high-rise' megacities in Asia and South America the average level of
234 buildings has continued to rise relative to the narrow spaces between buildings. In both types
235 of megacities, as L extends over 30-100 km, it becomes comparable with the size of L_R
236 mesoscale weather patterns (Hunt et al., 2004). Over this distance, winds tend to change
237 direction as air passes over the city and continues to affect the atmospheric boundary layer and
238 patterns of precipitation downwind (Cheng and Chan, 2012; Li et al., 2013a).

239

240 In low-rise cities whose forms, with few exceptions, are not changing significantly, the
241 populations are increasing, approximately in proportion to the area, i.e. L^2 (Hunt et al., 2011).
242 But in high-rise cities with rising population densities, the total population is rising at a faster
243 relative rate. Consequently, stationary energy sources such as those used in
244 heating/cooling/servicing buildings and for supplying industries and water (of particular
245 relevance in California; see Andrew, 2009) are increasing slightly more rapidly than L^2 .
246 However, the additional energy used for transportation (except in cities with high usage of
247 public transportation that are also more likely to integrate technological advancements – such
248 as those mentioned by Carrington (2016)) is increasing significantly more rapidly as cities
249 grow. Because the lengths of journeys in urban areas increase with L , and the population and
250 incomes are increasing as L^2 , the heat released per unit area by transportation is increasing in
251 proportion to L (Hunt et al., 2011). The energy for buildings, industry, and road transport is
252 generally supplied from outside built-up areas. New low carbon energy sources will therefore
253 be necessary to avoid the contribution of urban areas to global emissions of harmful pollution
254 (Kammen and Sunter, 2016).

255

256 The thermal environment affected by the city is determined by the heat capacity of buildings,
257 differing properties of urban land surfaces (e.g., absorption or reflection of solar radiation),
258 local heat emission from energy systems (e.g., air conditioning), and heat from transportation
259 (Sailor, 2011) and other anthropogenic sources such as re-radiation from buildings and traffic.
260 Equally important is the reduced ventilation caused by wind resistance of buildings, and in
261 some cities the reduction of solar radiation produced by dust and particles of air pollution.
262 These factors, which can be modified by planning, building design, and operation of urban

263 systems, alter the surface and temperatures inside and outside buildings (Georgescu et al.,
264 2015). During summer periods, temperatures are raised inside the urban area for longer periods
265 compared with rural areas, whose temperature decreases at a faster rate after sunset. The
266 average temperatures over 24 hours in urban areas can exceed rural temperatures by 5°C or
267 more, with serious implications on energy use and health, although this value has considerable
268 geographical and seasonal dependency. Importantly, the urban heat island (UHI) phenomenon
269 in high-density cities such as Hong Kong is rather due to anthropogenic impacts whereas in the
270 low-rise, less compact cities temperature increase is additionally governed by the re-emission
271 of energy absorbed by the built environment (Yang et al., 2016b). While the UHI can reduce
272 the need for heating in cold seasons, its effect will considerably increase energy use for cooling
273 purposes during the summer (see US Department of Energy, 2013). Compounding the
274 aforementioned changes in the physical environment are implications of, for example, extreme
275 heat events in cities, with broad health consequences (Hoshiko et al., 2010; O'Neill and Ebi,
276 2009; Hajat et al., 2007&2014; Huynen et al., 2001). But, if streets are covered with trees, or
277 roofs with vegetation, the shade and evapotranspiration lower daytime peak urban temperatures
278 (Maggiotto et al., 2014; Middel et al., 2015; Georgescu, 2015; Li and Norford, 2016; Tan et
279 al., 2016; Yang et al., 2016c). In large southeastern Asian cities (e.g., Li et al., 2013b), the UHI
280 typically exceeds 2°C at night. By contrast, in large cities in continental climates (with
281 populations greater than 5 million people) during high pollution episodes in winter, the urban
282 temperatures can be lower compared to rural areas throughout the diurnal cycle.

283
284 The effect of heat release in the city also affects the variation of temperature as the air moves
285 across the city, with maximum values where high-rise buildings are concentrated in the centre
286 of the city and towards the downwind side. The larger the city, the greater this effect is. Over
287 the neighbourhood scale of 1-3 km, temperatures are raised or lowered by parks, rivers,
288 buildings and the presence of other urban forms (Declat-Barreto et al., 2013; Connors et al.,
289 2013). A distribution of smaller parks lowers the average temperature more than a few large
290 parks (Bohnenstengel et al., 2011) and reduces the impacts of heat waves as shown in mortality
291 statistics for New York City (Huang, 2013). Such considerations bring to light the significance
292 of urban design and form (Zhou et al., 2011; Connors et al., 2013), which necessitate discussion
293 within a broader urban sustainability framework than has been acknowledged to date
294 (Georgescu et al., 2015).

295 296 Example 1. Energy use in a desert urban area

297 An example of the complexity by which urban areas can modify environmental hazards is
298 associated with the heat emitted by air conditioning (AC) systems. Physics-based modelling
299 simulations accounting for the variation in people's behaviour have been confirmed by
300 variations in observed temperatures. The results quantify the amount of electricity used on
301 diurnal time scales during a number of extreme heat events (EHEs) in a rapidly urbanizing
302 semi-arid metropolitan area (Phoenix, AZ), indicating that cooling from AC contributes about
303 53% of the overall daily electricity requirements during these periods. Electricity consumption
304 peaked during late afternoon hours (roughly 3-6pm, locally), when the demand for AC
305 approached two-thirds of the total hourly demand (Salamanca et al., 2013).

306

307 The multilayer building energy modelling (BEM) system, dynamically coupled to an
308 atmospheric model, was designed to predict cooling/heating energy demand (i.e., the energy
309 demand associated with ambient meteorological conditions) and has been applied at city scale
310 for contemporary and future conditions associated with urban expansion (Salamanca et al.,
311 2014, 2015). However, the BEM model alone is not able to predict the total energy demand
312 because the human behaviour consumption element (i.e., the energy component that is not
313 associated with the meteorology and therefore depends entirely on human behaviour) needs to
314 be accounted for separately.

315 Salamanca et al. (2013) estimated the human behaviour consumption (i.e., base load), analysing
316 citywide observed monthly mean electric loads for a specific year. For the Phoenix
317 metropolitan area, minimum observed electric loads occurred during March and November,
318 coinciding with moderate environmental weather conditions. These two months were
319 considered the baseline months with negligible heating/cooling electric consumption. In this
320 way and based on observed data, the diurnal cycle of the human behaviour consumption was
321 computed, coupled with the meteorological component, and used to calculate both electricity
322 consumption and its contribution to the region's UHI (Salamanca et al., 2013, 2014, 2015).

323 With higher peak temperatures and longer hot periods anticipated in future summers, electrical
324 demand by AC systems will have to be met by energy plants and the electric grid (Huang and
325 Gurney, 2016). Reliable energy forecasting methods, such as the simulations described above,
326 will be needed for resource planning of rapidly growing urban areas, especially in the extreme
327 conditions of semi-arid environments. Complicating such situations is the positive impact on
328 air quality associated with the destabilization of the planetary boundary layer (due to heat
329 emission from ACs), which promotes night-time vertical mixing and underscore challenges of
330 urban adaptation (Georgescu, 2015, Sharma et al., 2016). Therefore, compensating effects on
331 thermal, air quality, and other indicators, underline the need for comprehensive markers that
332 characterize the totality of urban-induced effects.

333

334 From an energy perspective, AC use is greatest during the same periods of extremely high
335 temperatures that cause higher transmission losses and reduced thermal efficiencies at electric
336 generation facilities. During a 2006 heat wave, electric power transformers failed in Missouri
337 and New York, causing interruptions of the electric power supply. In addition, more than 2,000
338 distribution line transformers in California failed during a July 2006 heat wave, causing loss of
339 power to approximately 1.3 million customers. Research ascertaining the potential for
340 individual and institutional adaptive strategies to lessen impacts due to extreme heat, and in
341 particular, impacts on human health risk caused by blackouts, is necessary to establish support
342 tools aiding development of novel protocols for heat risk emergency response monitoring and
343 planning (Kuras et al., 2017). Thus, increased cooling demand may increase the occurrence of
344 peak loads coinciding with periods when generation efficiencies are lowest. Furthermore, the
345 effects of high temperatures may be exacerbated when wind speeds are low or night-time
346 temperatures are high, preventing transmission lines from cooling. This is a particular concern

347 because night-time temperatures have been increasing at a faster rate than daytime
348 temperatures (e.g., Georgescu et al., 2013).

349 Comparison with observational data has demonstrated that the physics-based modelling system
350 is an effective tool for assessing (indoor) urban cooling requirements, which involves
351 evaluating electricity consumption for different urban growth patterns and under extreme
352 summertime weather conditions. These studies will be crucial for development of reliable
353 projections on future cooling needs and environmental consequences of rapidly urbanizing
354 regions under various climate scenarios (Georgescu et al., 2012, 2014; Bartos and Chester,
355 2015) that strategically incorporate adaptation and mitigation strategies alleviating energy
356 demand (Georgescu et al., 2014; Salamanca et al., 2016).

357 Example 2. Asian and subtropical cities

358 These cities have shown how, when very low regional temperatures occur, temperatures can
359 become even lower in urban areas as result of air pollution or sand storms reducing solar
360 radiation. The provision of heating that compensates for the cooling results in higher air
361 pollution, subsequently exacerbating hazards associated with extreme low temperatures. The
362 main hazards associated with pollutants in urban areas also arise from high concentrations of
363 contaminants from industry, transport, and agriculture, as well as particulates arising from
364 natural sources (e.g., wind-blown sand or noxious gases from lakes) (Jacobson, 2012; Li et al.
365 2015, 2016a), and may cause serious health implications (Pope and Dockery, 2006).

366
367 As winds transport pollutants into an urban area, concentrations tend to increase in the
368 downwind direction. At high temperatures during summer months, especially in the tropics,
369 climatic variations can induce low winds, high temperatures, which may be raised further by
370 high emissions from static and moving sources, such as episodes in Athens in 1987 (Matzarakis
371 and Mayer, 1991) and Moscow in 2010 (Shaposhnikov et al., 2014), and others in Beijing and
372 Shanghai (Wang and Gong, 2010; Huang et al., 2010). Because road vehicles are the main
373 source of polluting gases and particles in urban areas, and because journey distances (especially
374 for low-rise cities) increase as cities expand, emissions of air pollutants per unit area also
375 increase in proportion to the diameter L . The transport of atmospheric boundary layer
376 pollutants leads to degradation of air quality downwind, over distances that in some
377 meteorological conditions can extend hundreds of km (Cheng and Chan, 2012).

378
379 As air pollutants are transported across the city, while some gases increase in concentration,
380 others such as NO_x undergo chemical transformation and are reduced in the centre while
381 increasing in the outer regions. Overall, the magnitude of the pollutant concentration increases
382 with L . With anthropogenic climate change as an additional forcing agent, the sources of
383 pollutant and heat increase in urban areas can be compounded by, for example, widespread use
384 of AC use in buildings and vehicles. Both hazards are worsened by lengthy periods of calm
385 conditions, which are projected to occur with increased frequency under a future climate
386 characterized by increased occurrence of synoptic-scale blocking (Cassou and Guilyardi, 2007;
387 Li et al., 2012).

388

389 Example 3. Amplification of hazards in urban areas

390 Natural hazards arising outside urban areas are changed significantly within them. In some
391 situations, different urban hazards act in combination. In the presence of high winds, including
392 Tropical Cyclones (TCs) and tornadoes, although the resistance of buildings reduces the mean
393 wind speeds over the urban areas compared with outside, locally, wind speeds can exceed rural
394 wind speeds in gaps between buildings, and turbulent gusts are amplified (Oke, 1987; Britter
395 and Hunt, 1979). In addition, because significant numbers of high-rise buildings are being built
396 in often highly populous coastal cities that are subject to impacts from TCs, there is concern
397 about the growing vulnerability of their inhabitants (McGranahan et al., 2007; Pielke Jr., 2007).
398 Although global climate change is increasing the average sea surface temperature and the
399 average tropopause height, there is not yet any conclusive statistical evidence about the
400 projected strength and frequency of TCs. However, there is evidence that trajectories of those
401 major TCs that reach land are changing and reaching lower latitudes than previously (e.g., in
402 northern Malaysia). In general, the resistance to the airflow caused by the built environment
403 tends to deflect onshore winds parallel to the coast while amplifying peak near-surface winds
404 (Chan and Chan, 2015; Hunt et al., 2004). As urban areas expand, this trend is likely to be
405 amplified, which may also reduce the onshore movement of TCs. Coastal agricultural regions,
406 either surrounding or within urban areas themselves, that rely on TC rainfall may be negatively
407 affected if precipitation is reduced sufficiently, potentially resulting in increased irrigation
408 demand to meet required yields.

409

410 Hydrological extremes in the form of drought and flood can be amplified in urban areas. The
411 return period of intense precipitation over short periods (100 mm/hr) in Asia has decreased
412 (e.g., from 37 to 18 years according to Hong Kong Observatory (2016); cf. Wong et al., (2011)
413 and Lee et al., (2010)). The peak intensity of rainfall is likely to occur in geographical areas
414 where the surface air flow converges, which can happen in mountains, but also within
415 megacities, which affect regional climates (Shepherd et al., 2011; Georgescu et al., 2012; Smith
416 et al., 2013; Holst et al., 2016; Benson-Lira et al., 2016, Chow et al., 2016). The prediction of
417 rainfall and flooding in the low lying and almost completely urbanised areas of the island of
418 Singapore is improving as a result of detailed computational models and a dense network of
419 real time data (Pereira et al., 2014, Chow et al., 2016). Deeper convection caused by climate
420 change effects on the troposphere makes such events more likely in future. Important impacts
421 are also evident below the land-surface atmosphere interface. Decreased precipitation and
422 increased evaporation associated with longer periods of droughts and high temperature
423 episodes are depleting underground reservoirs and natural aquifers. In India, reduced monsoon
424 rains are lowering water levels in some lakes and rivers. Water shortages tend to be exacerbated
425 both within and around expanding urban areas, particularly in Asia and Africa where some city
426 aquifers are now more than 30m below ground level (Morris et al., 2003).

427

428 Flooding hazards in urban areas are partly caused by more rapid run-off from distant ice and
429 snow covered mountains caused by global warming, or by agricultural practices such as
430 reducing tree cover, which has been found to correlate with vulnerability against flooding (see

431 Pauleit et al., 2005 for a UK example). The results are seen in overflowing rivers and water
432 courses, and in unconfined areas such as streets and fields. Many secondary effects occur in
433 urban areas such as landslides, weaker foundations, collapsing structures (e.g., the Boulder,
434 CO, USA, floods of 2013 resulted in excess of \$1 billion in property damage) and loss of land
435 into coastal seas, depending on local geography and infrastructure. In the Philippines, these
436 secondary effects are found to influence the overall movement of floodwaters and the extent of
437 danger to communities (NOAH, 2012). The Tropical Cyclone Haiyan in 2013, upon reaching
438 the southern Philippines, caused unusually large damage to buildings and trees, in large part
439 because, over the shallow coastal waters, wind stress drove a large surge lifting rocks from the
440 seabed, transporting them several kilometres inland. Importantly, urban design strategies that
441 include incorporation of man-made rivers, reservoirs, and planned flood areas, have been
442 shown to reduce local flooding hazards relative to surrounding areas, highlighting the
443 importance of engineered infrastructure resilience as a potential adaptive mechanism.

444

445 Other hydrological hazards occur with wind and earthquake induced surges and waves (or
446 tsunamis) onto coasts (Fernando, 2008). Such hazards flood urban areas along coasts and along
447 canals connected to the coasts (as occurred with Hurricane Ike; see Kennedy et al., 2011).
448 Arctic coastal communities are now at risk from tsunamis generated by seismic activity that
449 has until now been suppressed by sea-ice. Examination of the tsunami waves of 2004, 2010,
450 and 2011 in Asia and the Pacific have illustrated how such hazards are affected by similar
451 physical and natural changes of climate (Klettner et al., 2012). As the tsunami in March 2011
452 reached the East coast of Japan, significant variability in the wave amplitude was observed and
453 in the surge movement (backwards and forwards up the shore) before flooding urban areas and
454 the industrial plant at Fukushima.

455 **3. Discussion and Conclusions**

456 This paper examines how extreme natural and artificial hazards in the atmosphere,
457 hydrosphere, and on land, are becoming more severe and more frequent as global climate
458 changes. The review emphasises the changing patterns and greater spatial variability of these
459 hazards over different geographical and climatic regions. As a result, the observed trends and
460 patterns of hazards are diverging from those of past decades and centuries. Existing models
461 suggest that this divergence will grow in the future with more intense, longer-duration, and
462 more frequent extreme events. The extent of the increase and variability in climate-induced
463 hazards varies between regions and for each specific hazard (Table 1).

464

465 In growing mega-cities, these factors include their surface area extent, urban population and
466 growth rate, urban design/technology, socio-economic factors and overall societal behaviour.
467 In considering the major physical and natural causes of hazards, this paper also describes and
468 analyses their societal effects, and in future work, will examine impacts on policy development.
469 The concept of societal effects extends beyond the useful, but essentially passive, concept of
470 vulnerability. For example, communities have shown increasing capability to obtain and use
471 information in advance of and during hazards (e.g., Hondula and Krishnamurthy, 2014), and

472 are increasingly adept at moderating hazard impacts (e.g., reducing the magnitudes and social
473 impacts of floods in urban areas; see Lagmay, 2015). The recovery of communities following
474 hazards can reduce long-term impacts. Equally significant are the interactions between urban–
475 biosphere interactions surrounding and within megacities, which have vast effects on health
476 and the agriculture/forestry municipalities in Latin America and southeast Asia.

477
478 The diagram shown in Figure 3 illustrates how the ‘dynamical-systems’ methodology (e.g.
479 Wilson, 2000) facilitates a holistic overview (Smuts, 1926), and informs decisions about the
480 empirical or scientifically based interactions, the various factors that influence or control some
481 broadly connected collection of processes and organisations. Here, we are considering the links
482 between global and regional climate change and the processes and hazards that affect urban
483 hazards, impacts, and potential ameliorating policies (see Hunt (2009b) for complex
484 relationships between these). The review presented here includes an appraisal of the impact on
485 health, in particular the combination of temperature extremes and intense air pollution from
486 traffic, heating and AC use, and particulates entering cities from rural areas. We stress the value
487 of comprehensive policy development accounting for place-based variability (Table 1) and
488 therefore directly address compensating effects on thermal, air quality, and other indicators that
489 characterize the totality of urban-induced effects. Simultaneously, we acknowledge that
490 wedge-type approaches (e.g., Pacala and Socolow, 2004) can provide insight into optimizing
491 the efficacy of urban policies, favouring some strategies over others.

492
493 We assert that involving different techniques for data analysis and system modelling is more
494 appropriate for practical decision making than the purely reductionist approach that builds up
495 semi-empirical models and connects them to basic data of all the various factors (e.g. Hunt et
496 al., 2012). Utility of such methods should ensure increasingly pragmatic approaches to
497 planning the form(s), size(s), and overall future growth of built environments, as well as
498 development of appropriate policies for green infrastructure and societal behaviour that will
499 lower energy use. To achieve success, however, will require action that is in concert with
500 societally informed decision-makers, grounded on sound scientific achievements. Collectively,
501 these actions will determine the future environment of mega-cities.

502

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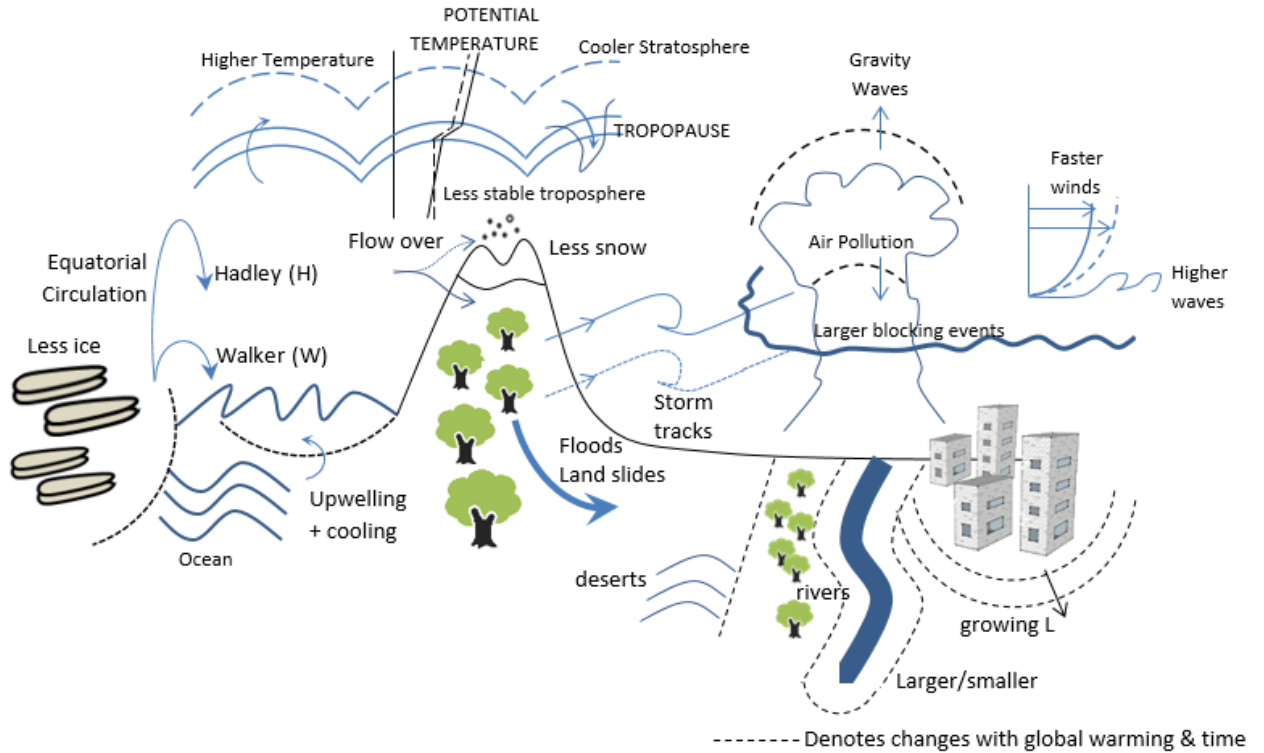
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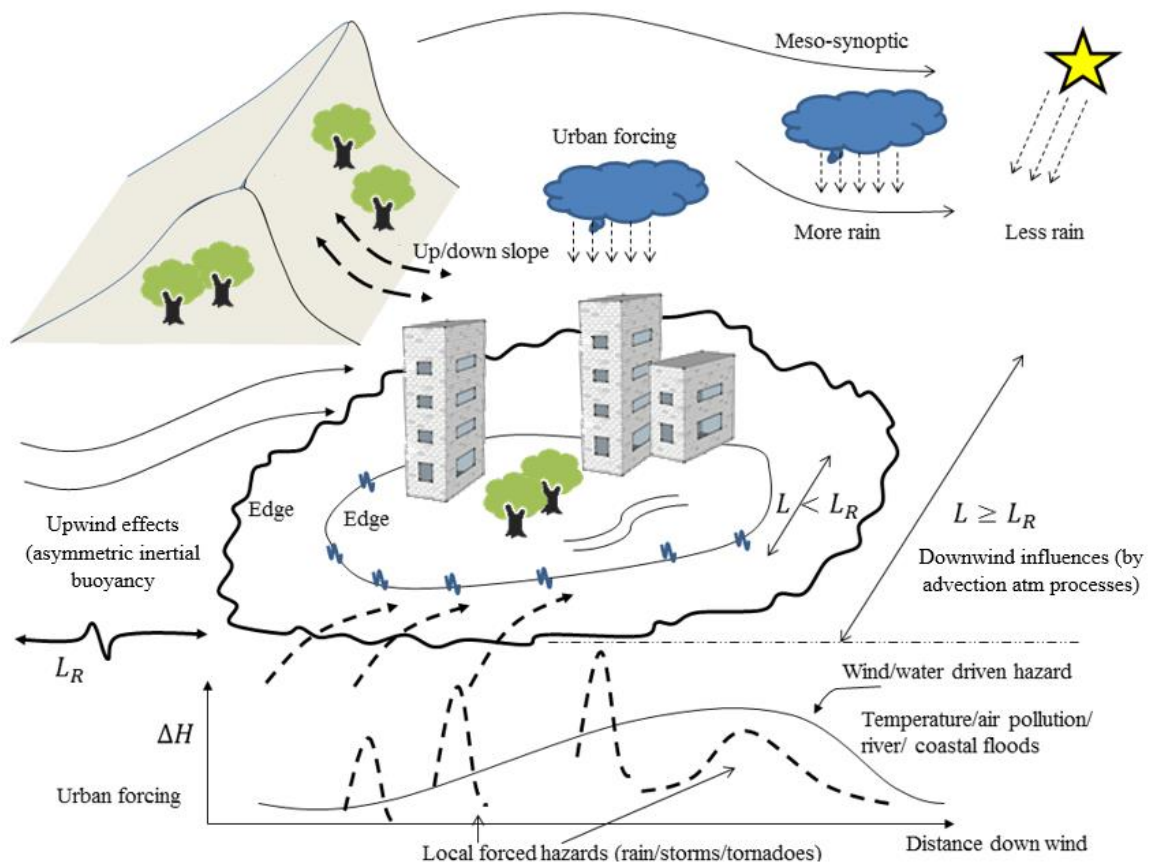
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906 **Figure 1:** Main natural effects and hazards influenced by climate change,

907 **,eg Atmosphere –changes in ocean-basin scale circulations ; vertical wind , temperature
908 and pollution profiles; Land –changes in urban ,rural, mountainous terrain ; Ocean and water-
909 changes in coupled ocean-atmosphere circulations; variable sea level rise; rivers grow or
910 shrink; generally less ocean and land ice



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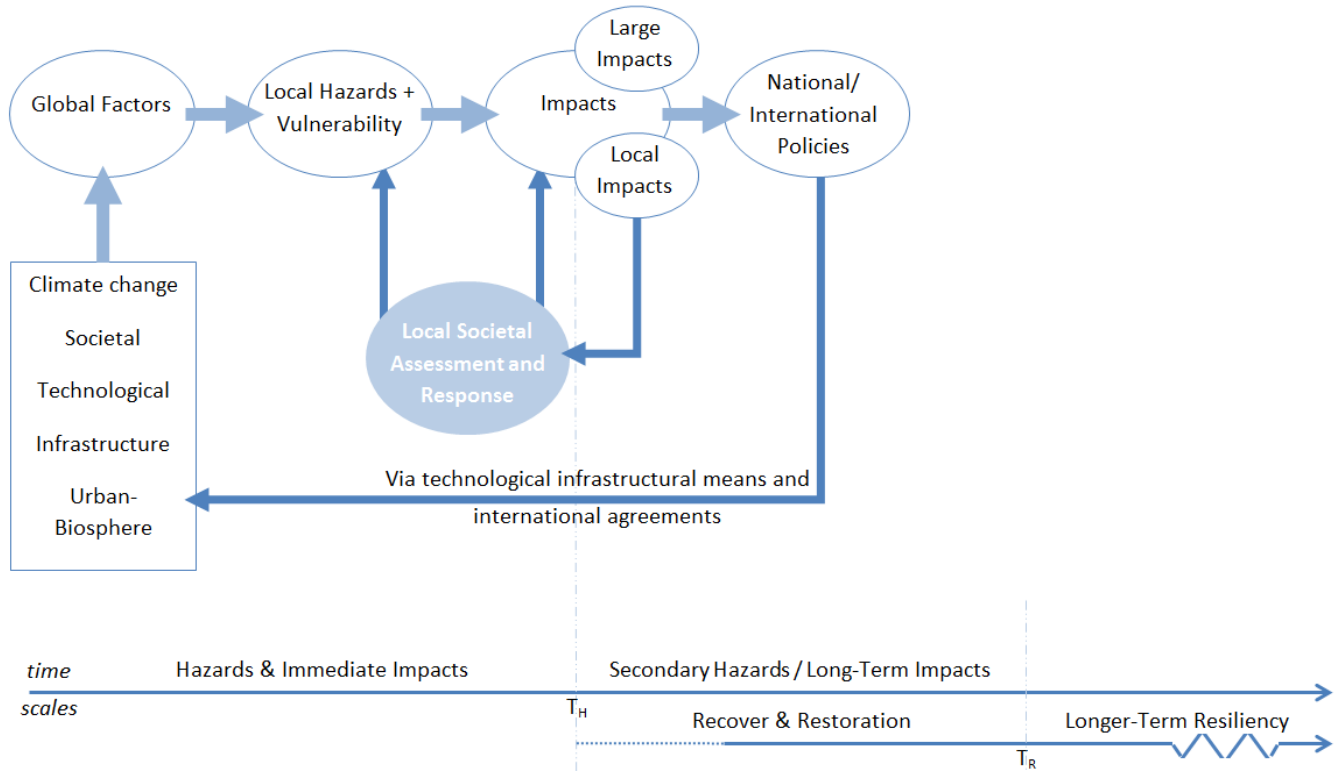
912 **Figure 2:** Increased hazards (ΔH) in urban areas in relation to urban forcing. When the size
 913 of the urban area (L) becomes equal to or greater than L_R (i.e., Rossby radius), the overall
 914 climatic hazards become typically more asymmetrical with a significant skew in the
 915 downwind direction. In this diagram air motion is from left to right, and, as shown in the
 916 graph at the bottom, air pollution, mean temperatures and cumulative wind driven hazard(s)
 917 increase with the distance downwind.

Table 1: Relating environmental hazards to urbanization and climate change factors.

<p>Types of natural and artificial hazards arising from extreme environmental variations (positive or negative denoted by +/-) and/or increased persistence and/or increased frequency in:</p>	<p>(i) wind speed (+/-)</p> <p>(ii) temperature (+/-)</p> <p>(iii) natural and artificial pollutants/radiation in the atmosphere, land, and water (+)</p> <p>(iv) hydrological processes, e.g. flooding, sea level rise/drought (+/-)</p> <p>(v) primary and secondary geophysical hazards leading to environmental hazards, e.g. earthquakes/tsunamis, volcanoes/land-slides, floods/land-slides, storm surges/floods, sand storms/air pollution (+)</p> <p>(vi) Biological/Environmental, e.g. disease, desertification (+)</p>
<p>Urbanisation and vulnerability effects</p>	<p>Amplification or reduction of natural hazards (listed above); hazards associated with infrastructure and human activities in urban areas (dependent on size, location, design and economy of urban areas); complex hazards associated with natural and human influences on global, regional, and urban environmental.</p> <p>e.g. higher/lower winds; higher temperatures, increased air pollution and increased/decreased water pollution; shortages/excesses of water.</p>
<p>Examples of significant hazards in large urban areas</p>	<p><u>Coastal and/or Riverine Cities</u> New Orleans (i), Houston (i), New York (i,v,vi), Bangkok (iv), Dhaka (i,iv), Tokyo (i, iii, v), Hong Kong (i, iii, iv), Jakarta (iii, v, vi), Manila (i, iii, iv), Paris and London (ii, iii, vi)</p> <p><u>Inland Cities</u> Phoenix (ii, iii), Beijing (ii, iii, vi) Xian/(other cities in central-western China)</p>

	(ii, iii, v), Moscow (ii, iii, vi), Athens (ii, iii, vi)
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920

921 **Figure 3:** System diagram for effects of hazards, impacts, and policies associated with global
 922 climate change, growing urban areas, and societal responses (T_H : period of hazards; T_R : period
 923 of long-term impacts and recovery).