

19 solutions are not significantly perturbed when open boundaries are placed close to the area
20 of interest. In more complex problems, this provides important performance improvements
21 in computational time as shown for a real application of harbor agitation.

22 **Keywords:** Open boundary, Perfectly Matched Layer, Mild-Slope equation, harbor.

23 INTRODUCTION

24 Harbor agitation by gravity waves is commonly predicted using elliptic Mild-Slope models
25 (Berkhoff 1972) in a semi-infinite domain, where the unbounded ocean part is connected to
26 physical boundaries (coastlines or harbor boundaries). Computational cost is compromised
27 by the size of the computational domain, directly related to the location of the artificial
28 or *open boundary*. This location is determined by the inherent hypothesis needed to define
29 the artificial boundary condition imposed and its accuracy. Thus, general and accurate
30 conditions generate smaller domains and provide important computational savings. The goal
31 of this paper is to propose an artificial boundary condition for coastal/harbor applications,
32 capable of reproducing the original semi-infinite solution at a minimum cost for the numerical
33 solver.

34 Standard strategies prescribing artificial non-reflecting boundary conditions impose an *a*
35 *priori* knowledge of the solution in the exterior domain. The artificial boundary condition
36 along the open boundary is constructed precisely as a consequence of the imposed solution
37 in the far-field area (e.g. Zubier et al. 2003; Li et al. 2005; Chen et al. 2005; Panchang
38 et al. 2008, among others). However, the exact solution in the outer semi-infinite domain
39 is, in general, not known a priori. Classical models exactly verify the Sommerfeld condition
40 through appropriate Hankel or Green functions (e.g. Tsay and Liu 1983; Xu and Panchang
41 1993), but for most real problems they are invalid because the classical models require to-
42 tally reflective and collinear coastlines. Xu et al. (1996) proposed a parabolic-approximation
43 based boundary condition to relax these limitations, but with the downside of not fulfilling
44 exactly the Sommerfeld condition. Moreover, this artificial open boundary requires a con-
45 stant bathymetry in the exterior domain. This is usually a strong assumption near the

46 coastline in semi-infinite domains. Panchang et al. (2000) addressed this issue by using
47 an incident wave that includes the exterior refractions provided by a non-constant idealized
48 bathymetry, resulting in important accuracy enhancements. Nevertheless, this strategy is
49 still affected by the parabolic-approximation based boundary condition and its limitations
50 (for instance, dominant radial direction of the scattered wave, and local bathymetry varia-
51 tions that are neglected on the open boundary). As a result, the open boundary requires
52 to be placed far enough from the region of interest, especially for complex shaped harbor
53 problems generating numerous reflections. This implies a larger computational domain and,
54 consequently, higher computational cost.

55 Here, a different approach overcoming the aforementioned limitations is proposed: the
56 Perfectly Matched Layer (PML) technique, originally proposed by Berenger (1994), is ex-
57 tended for the linear elliptic Mild-Slope including the exterior refractions provided by the
58 actual bathymetry. The PML method can be related to the dissipative *sponge layer* concept
59 applied in a number of coastal models, see for instance Larsen and Dancy 1983; Wei and
60 Kirby 1995; Lee and Yoon 2004; Sharma et al. 2014. The basic idea resides in surrounding
61 the interior domain by an artificial layer that aims to damp the diffracted wave energy of
62 the scattered wave. Both methods (sponge layer and PML) modify the original equations
63 inside the artificial layer. More precisely, the sponge layer method includes a decay reaction
64 term in the original Mild-Slope equation, see Sharma et al. (2014), while the derivatives
65 remain unchanged. Although this can dissipate the diffracted wave energy, it does not en-
66 sure a reflectionless interface connecting the interior domain with the artificial layer. As a
67 consequence, spurious reflections may pollute the solution. This issue can be alleviated by
68 enlarging the layer, but also increasing the computational cost.

69 On the contrary, the PML technique is constructed precisely to be a perfect reflectionless
70 artificial layer independently of the angle of incidence. PML transformation of the Mild-Slope
71 equation also affects the elliptic operator (derivatives), ensuring an analytical continuation
72 of the original solution inside the artificial layer (Teixeira and Chew 2000). Derivation of the

73 Mild-Slope PML equation following the primary rationale by Berenger (1994) is provided
74 in the Appendix. The PML has been applied in numerous scattering problems in the last
75 decade with excellent performance (e.g. Basu and Chopra 2003; Singer and Turkel 2004;
76 Michler et al. 2007; Demaldent and Imperiale 2013). Unbounded coastal problems with
77 PML and linearized shallow waters equations were addressed by Navon et al. (2004). In
78 harbor agitation with Mild-Slope models, the PML has been also recently used within a
79 model reduction framework by Modesto et al. (2015). These references for coastal modeling
80 are limited by constant exterior bathymetry, and hence no exterior refractions are considered.
81 First PML development with non-constant depth was addressed by Belibassakis et al. (2001),
82 but for complete unbounded domains and simple geometries.

83 In this paper, a strategy to incorporate the actual bathymetry is formulated for complex
84 harbor problems, using a rectangular shaped PML to approach the real semi-infinite do-
85 main. Here, the arbitrary definition of the bathymetry prevents an analytical continuation
86 of the solution in the PML, see Oskooi et al. (2008). In the aim of preserving the PML
87 reflectionless properties, some simplifications on the non-constant far-field bathymetry are
88 imposed. More precisely, this paper uses the idealization of the real water depth suggested by
89 Panchang et al. (2000), representing a good compromise between harbor models and reality.
90 Two different lateral sections of the idealized bathymetry are used, in order to retain the
91 exact water depth in the entire interior domain, and compute the incident wave. Moreover,
92 the proposed PML circumvents the difficulties associated to boundary conditions that apply
93 directly on semicircular artificial boundaries. For instance, it offers the possibility to incor-
94 porate fully non-collinear coastlines in a natural manner. Furthermore, it avoids the need
95 of interpolating the incident wave on the open boundary when two sections of the exterior
96 bathymetry are used (e.g. see Zhao et al. 2001). This results in satisfactory approximations
97 of the solution using very reduced computational domains, and for complex geometries that
98 generate numerous reflections.

99 The PML model is tested with four examples. Scattering on a circular object with

100 constant bathymetry is solved first to demonstrate the capabilities of the PML for very
 101 close open boundaries. A semicircular geometry with variable bathymetry is then used to
 102 verify that the idealized semi-infinite domain is properly reproduced. Validation through
 103 the standard elliptic shoal test for the Mild-Slope equation has been also addressed (briefly
 104 commented in the application section). Finally, to reinforce the validity of the model solution,
 105 an application to a real harbor located in the Northeast of Spain is shown.

106 ELLIPTIC HARBOR MODELS

107 In linear wave theory, the wave potential or complex surface elevation $\phi(x, y) \in \mathbb{C}$ prop-
 108 agates over the semi-infinite domain $\Omega_\infty \subset \mathbb{R}^2$ by means of the Mild-Slope equation

$$109 \quad \nabla \cdot (c c_g \nabla \phi) + k^2 c c_g \phi = 0 \quad \text{in } \Omega_\infty, \quad (1)$$

110 where $k(x, y) \in \mathbb{R}$ is the wavenumber, $c = \omega/k \in \mathbb{R}$ is the phase velocity, $\omega \in \mathbb{R}$ is the
 111 angular frequency of the monochromatic incident wave and $c_g \in \mathbb{R}$ is the group velocity. The
 112 wavenumber is related to frequency and to the slow varying bathymetry (i.e. mean-water-
 113 level-depth) $h(x, y) \in \mathbb{R}$ to account for refraction effects through the dispersion relation

$$114 \quad \omega^2 = kg \tanh(kh). \quad (2)$$

115 The group velocity is then defined as $c_g = d\omega/dk = g[\tanh(kh) + kh \operatorname{sech}^2(kh)]/(2\omega)$, with
 116 g the acceleration of gravity. Both Eqs. (1) and (2) can be properly modified to incorporate
 117 high bathymetry gradients and nonlinear effects, such as wave breaking or bottom friction
 118 (e.g. Booij 1981; Kirby 1984; Massel 1993).

119 This model requires boundary conditions everywhere on the boundary of the semi-infinite
 120 domain. Along coastlines and structures (breakwaters, walls, etc.) conforming the physical
 121 boundary, namely $\Gamma_{\mathbf{R}}$, the condition is given by Tsay and Liu (1983) as

$$122 \quad \mathbf{n} \cdot c c_g \nabla \phi - ikc c_g \alpha \phi = 0 \quad \text{on } \Gamma_{\mathbf{R}}, \quad (3)$$

123 where $i = \sqrt{-1}$ is the imaginary unit, \mathbf{n} is the outer unit normal at the boundary, and
 124 $\alpha \in [0, 1]$ is a real experimental coefficient controlling the reflection/absorption properties of
 125 the boundary. This coefficient is equal to zero on perfectly reflecting boundaries and to one
 126 on totally absorbing boundaries.

127 In addition, unbounded scattering problems require the Sommerfeld radiation condition
 128 that imposes that the scattered wave only has geometrical diffusion, namely

$$129 \quad \lim_{r \rightarrow \infty} \sqrt{r} \left(\frac{\partial}{\partial r} - ik \right) (\phi - \phi_0) = 0, \quad (4)$$

130 where r is the radial direction and $\phi_0(x, y) \in \mathbb{C}$ the imposed incident wave. Thus, $\phi - \phi_0$ is
 131 the scattered wave.

132 As commented before, Eq. (4) requires a special treatment in order to use a bounded
 133 computational domain and reproduce, on the artificial boundary, the effect of the Sommerfeld
 134 condition. The PML technique described next is a reasonable alternative in this case.

135 THE PML MODEL

136 It is standard practice to truncate the semi-infinite domain, Ω_∞ , into an interior region
 137 that includes the harbor/areas of interest, Ω_{int} , and a surrounding finite absorbing layer,
 138 Ω_{pml} . Thus, the computational domain, Ω , is the union $\bar{\Omega} = \bar{\Omega}_{\text{int}} \cup \bar{\Omega}_{\text{pml}}$, such that $\Omega \subset \Omega_\infty$.

139 The PML region is composed of four subdomains as shown in Figure 1, namely $\Omega_{\text{pml}} =$
 140 $\Omega_{\text{pml}}^{Lx} \cup \Omega_{\text{pml}}^{Rx} \cup \Omega_{\text{pml}}^y \cup \Omega_{\text{pml}}^{x,y}$. This PML is a rectangular shaped layer designed to absorb the
 141 scattered wave, along the Cartesian directions x and y , independently of its propagation
 142 angle.

143 The exterior bathymetry in both PML regions Ω_{pml}^{Lx} and Ω_{pml}^{Rx} is simplified accordingly to

144 the usual assumption for harbor models introduced by Panchang et al. (2000),

$$145 \quad h(x, y) = \begin{cases} h^L(y) & \text{if } (x, y) \in \Omega_{\text{pml}}^{Lx}, \\ h^R(y) & \text{if } (x, y) \in \Omega_{\text{pml}}^{Rx}, \\ h_0 & \text{if } (x, y) \in \Omega_{\text{pml}}^y \cup \Omega_{\text{pml}}^{x,y}. \end{cases} \quad (5)$$

146 This imposes an exterior bathymetry that varies only along the cross-shore direction as
 147 shown in Figure 1. The constant value h_0 corresponds to the far-field bathymetry region
 148 of the model. It determines the location of the PML at the top part of the computational
 149 domain. The exterior coastlines are modeled in the standard way, with straight lines parallel
 150 to the x axis, although the PML technique can be generally applied to arbitrary convex
 151 domains, see an example by Demaldent and Imperiale (2013). It is important to note that
 152 both non-constant exterior bathymetries, $h^L(y)$ and $h^R(y)$, are in general different in order
 153 to provide more faithful approximations of the unbounded domain. Moreover, the exterior
 154 coastlines have no collinearity assumptions in the present model, i.e. the coordinate y_c^L can
 155 be different from y_c^R as shown in Figure 1.

Under these assumptions, the mild-slope PML model reads

$$\nabla \cdot (c c_g \mathbf{P} \nabla \phi) + k^2 c c_g s_x s_y \phi = f(x, y) \quad \text{in } \Omega, \quad (6a)$$

$$\mathbf{n} \cdot (c c_g \mathbf{P} \nabla \phi) - i k c c_g \alpha \phi = 0 \quad \text{on } \Gamma_R, \quad (6b)$$

$$\mathbf{n} \cdot (c c_g \mathbf{P} \nabla \phi) - i k c c_g \phi = \mathbf{n} \cdot (c c_g \mathbf{P} \nabla \phi_0) - i k c c_g \phi_0 \quad \text{on } \Gamma_{\text{pml}}, \quad (6c)$$

156 where $\partial\Omega = \bar{\Gamma}_R \cup \bar{\Gamma}_{\text{pml}}$. Note that no Dirichlet boundary conditions are imposed. The
 157 non-homogeneous source term in (6a) is defined as

$$158 \quad f = \begin{cases} 0 & \text{if } (x, y) \in \Omega_{\text{int}}, \\ \nabla \cdot (c c_g \mathbf{P} \nabla \phi_0) + k^2 c c_g s_x s_y \phi_0 & \text{if } (x, y) \in \Omega_{\text{pml}}, \end{cases} \quad (7)$$

159 to account for the incident wave and to absorb only the scattered waves in the PML region.
 160 The diagonal anisotropy matrix \mathbf{P} defines the absorption in the PML area and it is the
 161 identity matrix in Ω_{int} , namely

$$162 \quad \mathbf{P} = \begin{pmatrix} s_y/s_x & 0 \\ 0 & s_x/s_y \end{pmatrix},$$

163 where $s_x = 1 + i\sigma_x/\omega$ and $s_y = 1 + i\sigma_y/\omega$ are two complex absorption parameters. The usual
 164 choice for the functions $\sigma_x(x)$ and $\sigma_y(y)$ are monotonic polynomials in the two respective
 165 Cartesian directions, namely

$$166 \quad \sigma_x(x) = \begin{cases} \sigma_0(|x - x_0|)^n/L_{\text{pml}} & \text{in } \Omega_{\text{pml}}^{Lx} \cup \Omega_{\text{pml}}^{Rx} \cup \Omega_{\text{pml}}^{x,y}, \\ 0 & \text{otherwise,} \end{cases} \quad (8a)$$

$$167 \quad \sigma_y(y) = \begin{cases} \sigma_0(y - y_0)^n/L_{\text{pml}} & \text{in } \Omega_{\text{pml}}^y \cup \Omega_{\text{pml}}^{x,y}, \\ 0 & \text{otherwise,} \end{cases} \quad (8b)$$

169 where the PML parameters are: σ_0 , absorption degree n and PML thickness L_{pml} . The
 170 interface $\Omega_{\text{int}} \cap \Omega_{\text{pml}}$ is assumed to be placed at coordinates x_0 and y_0 , see Figure 1. Note
 171 from this figure that the function σ_x is computed with $x_0 = x_0^L$ for Ω_{pml}^{Lx} , and with $x_0 = x_0^R$
 172 for Ω_{pml}^{Rx} . Note that in the interior domain the PML model does not change the original
 173 Mild-Slope Equations (MSE), recall Eq. (1), since $\sigma_x = \sigma_y = 0$ at any point $(x, y) \in \Omega_{\text{int}}$.

174 Eq. (6c) is a first-order *non-reflecting boundary condition* (Givoli 1992) on the PML outer
 175 boundary, and it is used to minimize spurious reflections. Other artificial conditions can be
 176 also used but, in practice, a proper choice of the absorbing parameters usually makes the
 177 solution not sensitive to the type of artificial boundary condition used on Γ_{pml} .

178 Finally, it is important to note that the PML model (6) behaves as a perfectly absorbing
 179 layer (i.e. it reproduces the original semi-infinite solution of Eq. (1)) only if the following

180 condition is satisfied: for any point in the PML region Ω_{pml} , the MSE (1) formulated in
 181 terms of the scattered wave must be homogeneous along the absorbing directions (x and y).
 182 That is, the scattered wave formulation particularized at the PML must have: (i) null source
 183 term, (ii) constant coefficients along the x direction in $\Omega_{\text{pml}}^{\text{Lx}} \cup \Omega_{\text{pml}}^{\text{Rx}} \cup \Omega_{\text{pml}}^{x,y}$, and (iii) constant
 184 coefficients along the y direction in $\Omega_{\text{pml}}^y \cup \Omega_{\text{pml}}^{x,y}$. Otherwise, the absorption properties of the
 185 artificial layer are not guaranteed, see Oskooi et al. (2008) and Oskooi and Johnson (2011)
 186 for details. It is straightforward to verify that these conditions are attained if:

187 (i) The incident wave ϕ_0 verifies the MSE (1) in the absorbing layer, that is

$$188 \quad \nabla \cdot (c c_g \nabla \phi_0) + k^2 c c_g \phi_0 = 0 \quad \text{in } \Omega_{\text{pml}}. \quad (9)$$

189 (ii) The exterior bathymetry satisfies Eq. (5).

190 Details on the PML derivation are shown in the Appendix. Next Section shows how the
 191 incident wave can be rapidly determined at any point $(x, y) \in \Omega_{\text{pml}}$ accordingly to Eq. (9).
 192 This follows the rationale proposed by Panchang et al. (2000).

193 **Computation of the incident wave field**

194 Note that Eq. (6) requires the expression of the incident wave only in the PML. In the
 195 far-field bathymetry region, the incident wave is defined as a monochromatic wave with an
 196 incoming angle $\theta \in \mathbb{R}$ and amplitude A_0 , namely

$$197 \quad \phi_0(x, y) = A_0 \exp(ik_0 x \cos \theta) \exp(ik_0 y \sin \theta) \quad \text{if } (x, y) \in \Omega_{\text{pml}}^y \cup \Omega_{\text{pml}}^{x,y}, \quad (10)$$

198 where k_0 is the wavenumber from (2) with $h = h_0$. Note that ϕ_0 , as defined by (10), verifies
 199 Eq. (9) with constant coefficients (i.e. the Helmholtz equation) in $\Omega_{\text{pml}}^y \cup \Omega_{\text{pml}}^{x,y}$. In the rest
 200 of the PML region the bathymetry depends only on coordinate y and is constant in the x
 201 (alongshore) direction. The procedure to compute ϕ_0 in this case consists in looking for an

202 incident wave with the same factorized structure as Eq. (10), that is

$$203 \quad \phi_0(x, y) = \exp(ik_0 x \cos \theta) \tilde{\phi}_0(y), \quad (11)$$

where $\tilde{\phi}_0$ is the cross-shore part of ϕ_0 . Substituting Eq. (11) in (9), and using proper boundary conditions, the function $\tilde{\phi}_0$ can be found as the solution of the second-order ordinary differential equation

$$\frac{\partial}{\partial y} \left(c c_g \frac{\partial \tilde{\phi}_0}{\partial y} \right) + \tilde{k}^2 c c_g \tilde{\phi}_0 = 0 \quad \text{in } y \in]y_c, y_0[, \quad (12a)$$

$$\tilde{\phi}_0(y_0) = A_0 \exp(ik_0 y_0 \sin \theta), \quad (12b)$$

$$\frac{\partial \tilde{\phi}_0}{\partial y}(y_c) + ik \tilde{\phi}_0(y_c) = 0, \quad (12c)$$

204 where $\tilde{k} = \sqrt{k^2 - k_0^2 \cos^2 \theta}$ is the associated wavenumber. Recall that parameters c , c_g and
 205 k depend only on coordinate y . The limits for the range of y , y_0 and y_c , correspond to
 206 the position of the far-field region and the coastline, respectively, see Figure 1. Note that
 207 boundary condition (12b) imposes continuity at the far-field region. On the other hand, Eq.
 208 (12c) corresponds to the one-dimensional version of the first-order artificial condition used
 209 in Eq. (6c). It imposes that the incident wave is not influenced by the harbor. No spurious
 210 reflections of $\tilde{\phi}_0$ are produced because the direction of incidence is obviously normal, and
 211 the first-order condition becomes in this case exact.

212 It is important to note that a distinguishing aspect of the proposed model is the possibility
 213 of using two different bathymetries in the regions $\Omega_{\text{pm1}}^{\text{Lx}}$ and $\Omega_{\text{pm1}}^{\text{Rx}}$, which can have also two
 214 different coastline positions, see Figure 1. This requires solving two times the ordinary
 215 differential equation (12): one using $y_c = y_c^{\text{L}}$ and the coefficients induced by the bathymetry
 216 $h^{\text{L}}(y)$, and another using $y_c = y_c^{\text{R}}$ and coefficients corresponding to $h^{\text{R}}(y)$. Note that a large
 217 distance between the two PML regions $\Omega_{\text{pm1}}^{\text{Lx}}$ and $\Omega_{\text{pm1}}^{\text{Rx}}$ reduces the impact of the idealization
 218 of the bathymetry, allowing better representations of the real semi-infinite domain. This

219 is a more realistic approximation than the standard used in elliptic harbors models, in
220 which either a unique cross-shore variation represents the whole exterior bathymetry, or
221 interpolation is required at the open boundary, see for instance Zhao et al. (2001) and Chen
222 et al. (2005).

223 **VALIDATION OF THE PML MODEL**

224 Three academic tests and one realistic example are considered next. The quantity of
225 interest used in all the examples is the amplification factor (i.e. normalized wave height),
226 namely

$$227 \quad H(x, y) = |\phi|/A_0. \quad (13)$$

228 The parameters of the PML have to be specified first. Recall that these parameters are
229 required for the definition of the absorbing functions in Eq. (8). Here, absorption degree
230 $n = 2$ is always used and the PML thickness L_{pml} is selected as 1.5 times the maximum wave
231 length induced by the lower frequency in each example, see Michler et al. (2007). The last
232 PML parameter, σ_0 , is set to $\sigma_0 = 60$ following the values proposed by Collino and Monk
233 (1998) that maximize the damping in the absorbing medium.

234 In general, the following examples use fourth-order triangular finite element meshes
235 adapted to the bathymetry, in order to achieve 8 nodes per wavelength over all the computa-
236 tional domain. Fourth-order elements are employed because of their numerical performance,
237 as demonstrated in the comparison study by Giorgiani et al. (2013). The complex linear
238 system resulted from this discretization is solved by means of direct methods.

239 **Circular scattering with constant bathymetry**

240 A standard test for open boundaries is first studied. The geometry is a reflecting circular
241 obstacle of radius R_c located with a distance R_d from the upper PML domain, as shown in
242 the left panel of Figure 2. These tests are standard to study the accuracy of the artificial
243 layers. The example consists in the scattering of an incident wave of period $T = 2\pi/\omega = 10$
244 s that propagates with an incoming angle $\theta = 0$. In this example the bathymetry is constant

245 in all the computational domain with value $h = 15.0119$ m from Panchang et al. 2000. The
 246 analytical solution for this problem is available in Mei (1983).

247 The computed amplification factor is depicted in Figure 2. Particularly, the left panel
 248 shows the case $R_d/R_c = 10$ and the right panel considers $R_d/R_c = 2$ with a not centered
 249 circle very close to the absorbing region. Note that the proximity of the PML does not
 250 perturb the symmetry of the solution. The relative error map between numerical and exact
 251 amplification factors (i.e. H and H^{exact} resp.) is depicted in Figure 3 for both cases, that
 252 is $|H - H^{\text{exact}}|/H^{\text{exact}}$, see Eq. (13). The maximum errors are on the order of 10^{-4} , that
 253 indicates an excellent agreement with the analytical solution even with a very close PML
 254 region.

255 This result outperforms previous harbor models based on the exterior description of the
 256 scattered wave, see for instance Panchang et al. (2000). Moreover, in order to be more
 257 confident about the reliability of the model, the convergence of the solution along the circle
 258 boundary (Γ_R) is shown in Figure 4. It is constructed by measuring the \mathcal{L}^2 norm of the
 259 relative error between approximation (ϕ) and exact solution (ϕ^{exact}) for different mesh sizes,
 260 namely

$$261 \quad E^2 = \frac{\int_{\Gamma_R} (\phi - \phi^{\text{exact}}) \overline{(\phi - \phi^{\text{exact}})} d\Gamma}{\int_{\Gamma_R} \phi^{\text{exact}} \overline{\phi^{\text{exact}}} d\Gamma}, \quad (14)$$

262 where the overline denotes the complex conjugate. High fidelity predictions (error level
 263 around 10^{-6}) are achieved, and the convergence rate for fourth-order finite elements is well
 264 reproduced in both tests (Babuška and Suri 1987). This corroborates the excellent behavior
 265 of the PML open boundary for constant bathymetry, even when it is placed in the vicinity
 266 of the obstacle.

267 **Semicircular scattering with variable bathymetry**

268 A scattering problem in a semi-infinite domain with variable bathymetry is explored next.
 269 The objective of this test is to show the ability of the proposed formulation to cope with two
 270 different bathymetries at the left and right boundaries. Furthermore, it is used to analyze

271 how the distance to the artificial boundary influences the results in an area of interest.

272 The geometry consists in a totally reflective boundary (i.e. $\alpha = 0$), including a semicircle
273 of radius R_c , which is adjacent to an absorbing boundary of length D_1 . The interior domain
274 is then limited by the vertical distance D_2 . Note from Figure 5 that distances D_1 and D_2
275 define the position of the PML in the geometry. The dimensionless variables $x' = x/R_c$,
276 $y' = y/R_c$, $D'_1 = D_1/R_c$ and $D'_2 = D_2/R_c$ are henceforth used. The bathymetry ranges as
277 $0.01 \leq h(x', y') \leq 0.3$ with $h(x', y') = 0.145y' + 0.053x' + 0.115$ in the region $-2 \leq x' \leq 2$,
278 and it becomes constant in the x' direction otherwise. This leads to two different cross-shore
279 bathymetries in the left and right regions, $x' < -2$ and $x' > 2$, respectively. Two oblique
280 incoming angles $\theta = \{220^\circ, 310^\circ\}$ and period $T = 0.6$ s are considered for testing the open
281 boundary.

282 The incident wave field is computed in the PML accordingly to Eq. (12) for each cross-
283 shore bathymetry. Exterior bathymetry effects are observed in the left region (see Figure
284 5), whereas the right cross-shore bathymetry does not induce noticeable refractions on the
285 far-field incident wave. Moreover, note from Figure 5 that this example induces strong
286 reflections in the distribution of the amplification factor. Obviously, these large values of the
287 wave amplification are inherent of this geometry (it presents a corner between two totally
288 reflective boundaries) and incoming wave direction. However, this does not influence the
289 purpose of this study because the unphysical amplification factors (i.e. larger than 3) are
290 independent of the PML application. In fact, this problem highlights the applicability and
291 excellent performance of the proposed PML method for problems with two different exterior
292 bathymetries.

293 In order to evaluate the influence of the PML in the solution of the problem, a set
294 of computations with the same mesh size but different combinations of D'_1 and D'_2 are
295 performed. Specifically, fourth order uniform meshes with a minimum resolution of 16 nodes
296 per wavelength are employed. Since there is no analytical solution for this problem, a
297 reference computation with $D'_1 = D'_2 = 4$ and half of the element size is used to evaluate

298 the error. On the semicircle boundary, named as Γ , the following errors in the amplification
 299 factor are defined: the mean error

$$300 \quad E_1^2 = \frac{1}{\text{meas}(\Gamma)} \int_{\Gamma} (H^* - H)^2 d\Gamma, \quad (15)$$

301 and the maximum elemental error, namely

$$302 \quad E_2^2 = \max_{\forall \Gamma_i \subset \Gamma} \frac{1}{\text{meas}(\Gamma_i)} \int_{\Gamma_i} (H^* - H)^2 d\Gamma, \quad (16)$$

303 where H is the computed (approximated) amplification factor, as defined in Eq. (13), and
 304 H^* is the reference solution. Each Γ_i is the side of a finite element along Γ .

305 These errors are depicted in Figure 6 for a variety of values of D'_1 and D'_2 . For both mean
 306 and maximum errors, the parameter D'_2 that stands for the PML position in the far-field
 307 (constant) bathymetry region has no influence on the solution along the semicircle boundary.
 308 The errors are measured on the semicircular scatterer, if the incoming wave angle is $\theta = 220^\circ$
 309 waves are only slightly refracted and D'_2 may seem not very influential. Thus, an incoming
 310 wave angle $\theta = 310^\circ$ is also tested and the errors plotted, see Figure 6 (right).

311 The results produced by the parameter D'_2 are in agreement to the conclusions of the
 312 previous example in which the bathymetry was also constant. In fact, the far-field PML
 313 region can be placed at a minimum (optimal) distance of $D'_2 = 0$. As expected, this behavior
 314 is not reproduced for the parameter D'_1 defining the PML position in the variable bathymetry
 315 regions. Nevertheless, it is observed that using only $D'_1 = 1$, and hence with the PML
 316 very close to the obstacle, is sufficient to ensure that even the maximum elemental error
 317 is no longer (significantly) perturbed. As seen in Figure 6, for $D'_1 > 1$ the error is almost
 318 constant and only due to the finite element discretization. It is important to remark that
 319 each dot in Figure 6 represents a finite element computation for a given value of D'_1 and
 320 D'_2 . These computation are done with finite element meshes with the same characteristic
 321 element size (same error bound) but not with identical meshes around the semi-circular

322 scatterer (the nodes are not exactly in the same position in the smaller interior domain
 323 defined by $D'_1 = D'_2 = 0$). This induces negligible differences that are more evident when the
 324 curve flattens, that is, when the PML error is negligible compared with the finite element
 325 one (the discretization error).

326 In order to obtain more information on the PML influence, Figure 7 depicts the accuracy
 327 of the computed amplification factor in the smaller interior domain (defined by $D'_1 = D'_2 = 0$).
 328 Accuracy is shown by the number of correct significant digits of the solution (see Remark
 329 1). The isolines of amplification factor equal to one ($H = 1$) for both the computed and
 330 reference solutions are also shown. In the three depicted computations, $D'_2 = 0$ is selected
 331 since this parameter has not significant influence on the results. Convergence to the reference
 332 solution when increasing the value of D'_1 is observed. Note that the area within the isoline of
 333 one correct digit of accuracy decreases as D'_1 increases. The case $D'_1 = 1$, that provides the
 334 optimal PML position when measuring the global error on the circle, produces an accuracy
 335 of one correct significant digit in almost all the area of interest. The regions with no accuracy
 336 are negligible and almost disappear as D'_1 increases. Consequently, in this synthetic problem
 337 with non-constant exterior (idealized) bathymetry, the PML is also able to reproduce the
 338 semi-infinite solution using a small computational domain.

339 **Remark 1** (Correct significant digits of the solution). Numerical accuracy can be evaluated
 340 by estimating the number of correct significant digits (or correct significant figures) between
 341 approximation (ϕ) and reference solution (ϕ^{ref}). An approximation has q significant digits
 342 if the relative error, namely $e = (\phi - \phi^{\text{ref}})/\phi^{\text{ref}}$, verifies $|e| < 0.5 \times 10^{-q}$, or alternatively,
 343 $q = -\log_{10} |e|$, see Higham (2002). That is, the first nonzero digit of the approximation and
 344 up to q succeeding digits can be “trusted”. From Higham (2002): “an approximation ϕ to
 345 ϕ^{ref} has q correct significant digits if ϕ and ϕ^{ref} round to the same number to q significant
 346 digits”. Note that applications for harbor agitation usually do not demand more than $q = 2$.

347 **Elliptic shoal**

348 The elliptic shoal is a well-known test case largely used to validate models in coastal
349 regions. It consists in testing the refraction effects produced by a shoal located into a mild-
350 sloped bathymetry plane. Although this example does not present large wave reflections,
351 it has been used here to validate the proposed methodology with reference experimental
352 and numerical results. The definition of the bathymetry and shoal limits can be found, for
353 instance, in Belibassakis et al. (2001).

354 The results of this test (not shown) were nearly identical to those in Oliveira and Anasta-
355 siou (1998) and Panchang et al. (1991). That is, it performs fairly well within the limitations
356 of the linear wave theory, that usually over-predicts the energy concentration associated to
357 the shoal refraction. Nonlinear wave models, see for instance Woo and Liu (2004), can be
358 used to drastically improve the wave description in this particular test.

359 **Mataró Harbor**

360 This example is presented to show the applicability and reliability of the proposed model
361 for more complex problems. The real geometry and bathymetry of the Mataró harbor,
362 located in the Northeast of Spain, are considered. Figure 8 shows three computational
363 domains used for comparison purposes. The small domain hardly includes the exterior
364 part of the harbor and it produces a 80% of reduction in degrees of freedom (DOF) with
365 respect to the largest domain, and 40% with respect to the medium domain. The medium
366 computational domain (60 000 DOF) has a 70% of reduction with respect to the largest
367 domain and is probably a reasonable engineering option. Note that the largest domain is
368 used to obtain a reference solution for this problem, and it incorporates the real bathymetry
369 and coastlines. Nevertheless, both small and medium domains are drastically smaller than
370 the reference one and they include two different exterior bathymetries with completely non-
371 collinear coastlines. Note also from Figure 8 that models requiring collinear coastlines cannot
372 reproduce the real geometry using small domains. In terms of the computational time
373 required to solve this problem (assembling of matrices and solution of linear systems), the

374 medium and small domains, respectively, use the 29% and 17% of the CPU time demanded
375 by the largest domain.

376 The absorbing coefficient α in Eq. (3) is specified in Figure 8 for all the boundaries of the
377 medium computational domain. The rest of boundaries correspond to the case $\alpha = 0.7$. For
378 each domain, both short (6 s) and longer (16 s) incident waves are used in the computations
379 accordingly to those limit values observed offshore in the region. Two different incoming
380 directions, $\theta = 270^\circ$ and $\theta = 225^\circ$, are also considered to cover normal and oblique incident
381 waves. A particular solution of the amplification factor for the shorter and oblique incident
382 waves is depicted in Figure 8.

383 In order to test the model performance, three different sections of interest are defined in
384 the harbor (see the medium computational domain in Figure 8). Results of the amplification
385 factor on these sections, for shorter and longer waves, are depicted in Figures 9 and 10,
386 respectively. In each figure, the top row depicts the comparison between different domains
387 for the normal incoming wave direction, whereas the bottom row shows the oblique case.

388 In general, the small domain produces reasonably good results, under engineering toler-
389 ance, following the trend of the reference solution specially for the wave phase. This indicates
390 that PML is able to reproduce the exterior effects, even though it only uses a very small
391 region of the exterior part of the harbor. It is also observed that the medium domain re-
392 duces the error in almost all the cases. The minor differences are observed for the shorter
393 and normal incident waves (top row in Figure 9), because under these circumstances the
394 influence of the exterior part of the harbor decreases considerably on the sections of interest.
395 As expected, the oblique incoming direction induces larger errors (bottom row in Figure 9).
396 In general, the smaller domain slightly underestimates the wave amplification. Note that
397 the medium domain produces very satisfactory results even for the oblique case.

398 The longer waves case is specially interesting because the exterior bathymetry produces
399 more refractions than for short waves. A comparison is depicted in Figure 11 for the com-
400 puted incident wave. In previous models based on the exterior description of the scattered

401 wave, these refractions induce higher errors in the computations associated to large wave pe-
402 riods, see Panchang et al. (2000). Here, the results from Figure 10 show that the differences
403 with respect to the reference solution are similar to the short waves case. This indicates less
404 sensitivity of the PML model to variations on the wave period.

405 **CONCLUDING REMARKS**

406 The Perfectly Matched Layer method is proposed to model an artificial non-reflecting
407 boundary for elliptic harbor models. A rectangular shaped absorbing layer is introduced,
408 and the linear Mild-Slope equation is used to describe the wave physics in the harbor. Ex-
409 terior refractions are incorporated into the model by using a natural simplification of the
410 exterior domain, geometry and bathymetry. Particularly, the exterior bathymetry is con-
411 sidered constant only in the alongshore direction. As a result of this not very restrictive
412 assumption, the incident wave is efficiently computed in the absorbing region as the solu-
413 tion of a one-dimensional problem. Moreover, and in opposite to standard strategies, two
414 exterior bathymetries are used with two non-collinear coastlines, in order to more accurately
415 approximate the real unbounded domain.

416 Results show that solutions are not significantly perturbed even with close placed artificial
417 boundaries. This is specially evident in the far-field constant bathymetry area of the model.
418 In this case, the top part of the PML is placed practically on the limit of constant bathymetry.

419 The proposed model is promising also for more complex problems. For a real case, the
420 model produces good results with a computational domain that hardly involves the exterior
421 part of the harbor. This is observed even with an oblique incoming wave direction and a
422 large period. Reliable solutions of the harbor agitation are therefore obtained at a reduced
423 computational cost. Future advancements include the use of the PML in more sophisticated
424 propagation models for harbor agitation, including non-linear wave interactions.

425 **APPENDIX I. DERIVATION OF THE PML**

426 This appendix describes the derivation of Eq. (6a) in the PML. The original rationale
427 from Berenger (1994) imposes the absorption of the scattered wave through the transient

428 wave propagation solution, although other methods can be used to obtain the same equation,
 429 see Chew et al. (1997). The original rationale is followed here.

430 In the PML, the transient Mild-Slope equation, where $\tilde{\phi}_s(x, y, t) = \phi_s(x, y)e^{-i\omega t}$ is the
 431 monochromatic scattered wave and $\phi_s = \phi - \phi_0$, can be written as

$$432 \quad \frac{\partial^2 \tilde{\phi}_s}{\partial t^2} - \frac{c}{c_g} \nabla \cdot (cc_g \nabla \tilde{\phi}_s) = 0 \quad \text{in } \Omega_{\text{pml}}. \quad (17)$$

433 Note that null RHS is imposed because the incident wave fulfills the equation in the PML,
 434 recall Eq. (9). The change of variables

$$435 \quad \tilde{v} = \frac{\partial \tilde{\phi}_s}{\partial t}, \quad \tilde{q}_x = \frac{\partial \tilde{\phi}_s}{\partial x}, \quad \tilde{q}_y = \frac{\partial \tilde{\phi}_s}{\partial y}, \quad (18)$$

and the decomposition $\tilde{v} = \tilde{v}_x + \tilde{v}_y$ lead to the first order system

$$\frac{\partial \tilde{v}_x}{\partial t} = \frac{c}{c_g} \frac{\partial}{\partial x} (cc_g \tilde{q}_x), \quad (19a)$$

$$\frac{\partial \tilde{q}_x}{\partial t} = \frac{\partial \tilde{v}_x}{\partial x} + \frac{\partial \tilde{v}_y}{\partial x}, \quad (19b)$$

$$\frac{\partial \tilde{v}_y}{\partial t} = \frac{c}{c_g} \frac{\partial}{\partial y} (cc_g \tilde{q}_y), \quad (19c)$$

$$\frac{\partial \tilde{q}_y}{\partial t} = \frac{\partial \tilde{v}_x}{\partial y} + \frac{\partial \tilde{v}_y}{\partial y}. \quad (19d)$$

This system provides a solution for Eq. (17). Note that Eqs. (19a) and (19b) describe the propagation of \tilde{v}_x and \tilde{q}_x , respectively, along the x direction. Analogously, Eqs. (19c) and (19d) describe the propagation of \tilde{v}_y and \tilde{q}_y along the y direction. Note, moreover, that Eqs. (19a) and (19c) are fully homogeneous (constant cc_g) in the PML along their directions because of the bathymetry restriction (5). The main point of the PML resides in modifying the original system (19) with a damping term in each direction that vanishes outside the

PML, that is

$$\frac{\partial \tilde{v}_x}{\partial t} + \sigma_x \tilde{v}_x = \frac{c}{c_g} \frac{\partial}{\partial x} (cc_g \tilde{q}_x), \quad (20a)$$

$$\frac{\partial \tilde{q}_x}{\partial t} + \sigma_x \tilde{q}_x = \frac{\partial \tilde{v}_x}{\partial x} + \frac{\partial \tilde{v}_y}{\partial x}, \quad (20b)$$

$$\frac{\partial \tilde{v}_y}{\partial t} + \sigma_y \tilde{v}_y = \frac{c}{c_g} \frac{\partial}{\partial y} (cc_g \tilde{q}_y), \quad (20c)$$

$$\frac{\partial \tilde{q}_y}{\partial t} + \sigma_y \tilde{q}_y = \frac{\partial \tilde{v}_x}{\partial y} + \frac{\partial \tilde{v}_y}{\partial y}, \quad (20d)$$

where $\sigma_x(x)$ and $\sigma_y(y)$ are defined in Eq. (8). The frequency domain version of this system considers a monochromatic definition of the unknowns, namely $\tilde{v}(x, y, t) = v(x, y)e^{-i\omega t}$, $\tilde{q}_x(x, y, t) = q_x(x, y)e^{-i\omega t}$ and $\tilde{q}_y(x, y, t) = q_y(x, y)e^{-i\omega t}$. System (20) is then rewritten as

$$(-i\omega + \sigma_x)v_x = \frac{c}{c_g} \frac{\partial}{\partial x} (cc_g q_x),$$

$$-i\omega q_x = \frac{1}{s_x} \left(\frac{\partial v_x}{\partial x} + \frac{\partial v_y}{\partial x} \right),$$

$$(-i\omega + \sigma_y)v_y = \frac{c}{c_g} \frac{\partial}{\partial y} (cc_g q_y),$$

$$-i\omega q_y = \frac{1}{s_y} \left(\frac{\partial v_x}{\partial y} + \frac{\partial v_y}{\partial y} \right),$$

436 where $s_x = 1 + i\sigma_x/\omega$ and $s_y = 1 + i\sigma_y/\omega$. After rearranging all these equations and using
 437 the decomposition $v = v_x + v_y$ and the change of variables (18), the following equation for
 438 the scattered wave $\phi_s(x, y)$ arises

$$439 \quad -\frac{c_g}{c} \omega^2 \phi_s = \frac{1}{s_x} \left(\frac{cc_g}{s_x} \frac{\partial \phi_s}{\partial x} \right) + \frac{1}{s_y} \left(\frac{cc_g}{s_y} \frac{\partial \phi_s}{\partial y} \right) \quad \text{in } \Omega_{\text{pml}}. \quad (22)$$

440 Taking into account that $k = \omega/c$ and $\phi = \phi_s + \phi_0$, the Mild-Slope equation (6a) is obtained.

441 ACKNOWLEDGMENTS

442 This work has been partially supported by the Spanish Ministry of Science and Competi-
 443 tiveness, grants number: DPI2011-27778-C02-02 and DPI2014-51844-C2-02, and the Catalan

444 Government, grant number: 2009SGR875.

445 REFERENCES

- 446 Babuška, I. and Suri, M. (1987). “The optimal convergence rate of the p-version of the finite
447 element method.” *SIAM J. Numer. Anal.*, 24(4), 750–776.
- 448 Basu, U. and Chopra, A. K. (2003). “Perfectly matched layers for time-harmonic elasto-
449 dynamics of unbounded domains: Theory and finite-element implementation.” *Comput.*
450 *Methods Appl. Mech. Eng.*, 192(11-12), 1337–1375.
- 451 Belibassakis, K. A., Athanassoulis, G. A., and Gerostathis, T. P. (2001). “A coupled-
452 mode model for the refraction-diffraction of linear waves over steep three-dimensional
453 bathymetry.” *Appl. Ocean Res.*, 23(6), 319–336.
- 454 Berenger, J.-P. (1994). “A perfectly matched layer for the absorption of electromagnetic
455 waves.” *J. Comput. Phys.*, 114(2), 185–200.
- 456 Berkhoff, J. C. W. (1972). “Computation of combined refraction-diffraction.” *Proc. 13th*
457 *Coastal Engineering Conference*, Vol. 1, Vancouver, Canada. 471–490.
- 458 Booij, N. (1981). “Gravity waves on water with non-uniform depth and current.” *Report No.*
459 *81-1*, Department of Civil Engineering, Delft University of Technology, The Netherlands.
- 460 Chen, W., Panchang, V., and Demirbilek, Z. (2005). “On the modeling of wave-current
461 interaction using the elliptic mild-slope wave equation.” *Ocean Eng.*, 32(17-18), 2135–
462 2164.
- 463 Chew, W. C., Jin, J. M., and Michielssen, E. (1997). “Complex coordinate stretching as a
464 generalized absorbing boundary condition.” *Microw. Opt. Technol. Lett.*, 15(6), 363–369.
- 465 Collino, F. and Monk, P. (1998). “Optimizing the perfectly matched layer.” *Comput. Methods*
466 *Appl. Mech. Eng.*, 164(1–2), 157–171.
- 467 Demaldent, E. and Imperiale, S. (2013). “Perfectly matched transmission problem with
468 absorbing layers: Application to anisotropic acoustics in convex polygonal domains.” *Int.*
469 *J. Numer. Methods Eng.*, 96(11), 689–711.
- 470 Giorgiani, G., Modesto, D., Fernández-Méndez, S., and Huerta, A. (2013). “High-order con-

471 tinuous and discontinuous Galerkin methods for wave problems.” *Int. J. Numer. Methods*
472 *Fluids*, 73(10), 883–903.

473 Givoli, D. (1992). *Numerical methods for problems in infinite domains*, Vol. 33 of *Studies in*
474 *Applied Mechanics*. Elsevier Scientific Publishing Co., Amsterdam.

475 Higham, N. J. (2002). *Accuracy and Stability of Numerical Algorithms*. Society for Industrial
476 and Applied Mathematics, Philadelphia, PA, USA.

477 Kirby, J. T. (1984). “Note on linear surface wave-current interaction over slowly varying
478 topography.” *J. Geophys. Res.*, 89(NC1), 745–747.

479 Larsen, J. and Dancy, H. (1983). “Open boundaries in short wave simulations - A new
480 approach.” *Coast. Eng.*, 7(3), 285–297.

481 Lee, C. and Yoon, S. B. (2004). “Effect of higher-order bottom variation terms on the
482 refraction of water waves in the extended mild-slope equations.” *Ocean Eng.*, 31(7), 865–
483 882.

484 Li, D., Panchang, V., Tang, Z., Demirbilek, Z., and Ramsden, J. (2005). “Evaluation of an
485 approximate method for incorporating floating docks in harbor wave prediction models.”
486 *Can. J. Civ. Eng.*, 32(6), 1082–1092.

487 Massel, S. R. (1993). “Extended refraction-diffraction equation for surface waves.” *Coast.*
488 *Eng.*, 19(1-2), 97–126.

489 Mei, C. C. (1983). *The applied dynamics of ocean surface waves*. Wiley, New York.

490 Michler, C., Demkowicz, L., Kurtz, J., and Pardo, D. (2007). “Improving the performance
491 of perfectly matched layers by means of hp-adaptivity.” *Numer. Meth. Part. Differ. Equ.*,
492 23(4), 832–858.

493 Modesto, D., Zlotnik, S., and Huerta, A. (2015). “Proper generalized decomposition for
494 parameterized Helmholtz problems in heterogeneous and unbounded domains: Application
495 to harbor agitation.” *Comput. Methods Appl. Mech. Eng.*, 295, 127–149.

496 Navon, I. M., Neta, B., and Hussaini, M. Y. (2004). “A perfectly matched layer approach to
497 the linearized shallow water equations models.” *Mon. Wea. Rev.*, 132(6), 1369–1378.

498 Oliveira, F. S. B. F. and Anastasiou, K. (1998). “An efficient computational model for water
499 wave propagation in coastal regions.” *Appl. Ocean Res.*, 20(5), 263–271.

500 Oskooi, A. and Johnson, S. G. (2011). “Distinguishing correct from incorrect PML proposals
501 and a corrected unsplit PML for anisotropic, dispersive media.” *J. Comput. Phys.*, 230(7),
502 2369–2377.

503 Oskooi, A. F., Zhang, L., Avniel, Y., and Johnson, S. G. (2008). “The failure of perfectly
504 matched layers, and towards their redemption by adiabatic absorbers.” *Opt. Exp.*, 16(15),
505 11376–11392.

506 Panchang, V., Chen, W., Xu, B., Schlenker, K., Demirbilek, Z., and Okihiro, M. (2000).
507 “Exterior bathymetric effects in elliptic harbor wave models.” *J. Waterw. Port Coast.*
508 *Ocean Eng.*, 126(2), 71–78.

509 Panchang, V., Pearce, B., Wei, G., and Cushman-Roisin, B. (1991). “Solution of the mild-
510 slope wave problem by iteration.” *Appl. Ocean Res.*, 13(4), 187–199.

511 Panchang, V., Zhang, J., and Demirbilek, Z. (2008). “Incorporating rubble mound jetties in
512 elliptic harbor wave models.” *J. Waterw. Port Coast. Ocean Eng.*, 134(1), 40–52.

513 Sharma, A., Panchang, V., and Kaihatu, J. (2014). “Modeling nonlinear wavewave interac-
514 tions with the elliptic mild slope equation.” *Appl. Ocean Res.*, 48, 114–125.

515 Singer, I. and Turkel, E. (2004). “A perfectly matched layer for the Helmholtz equation in a
516 semi-infinite strip.” *J. Comput. Phys.*, 201(2), 439–465.

517 Teixeira, F. L. and Chew, W. C. (2000). “Complex space approach to perfectly matched
518 layers: a review and some new developments.” *Int. J. Numer. Model. Electron. Networks*,
519 13(5), 441–455.

520 Tsay, T.-K. and Liu, P. L.-F. (1983). “A finite element model for wave refraction and diffrac-
521 tion.” *Appl. Ocean Res.*, 5(1), 30–37.

522 Wei, G. and Kirby, J. T. (1995). “Time-dependent numerical code for extended Boussinesq
523 equations.” *J. Waterw. Port Coast. Ocean Eng.*, 121(5), 251–261.

524 Woo, S.-B. and Liu, P. L.-F. (2004). “Finite-element model for modified Boussinesq equa-

525 tions. I: Model development.” *J. Waterw. Port Coast. Ocean Eng.*, 130(1), 1–16.

526 Xu, B. and Panchang, V. (1993). “Outgoing Boundary Conditions for Finite-Difference El-
527 liptic Water-Wave Models.” *Proc. R. Soc. London Ser. A*, 441(1913), 575–588.

528 Xu, B., Panchang, V., and Demirbilek, Z. (1996). “Exterior reflections in elliptic harbor
529 wave models.” *J. Waterw. Port Coast. Ocean Eng.*, 122(3), 118–126.

530 Zhao, L., Panchang, V., Chen, W., Demirbilek, Z., and Chhabbra, N. (2001). “Simulation of
531 wave breaking effects in two-dimensional elliptic harbor wave models.” *Coast. Eng.*, 42(4),
532 359–373.

533 Zubier, K., Panchang, V., and Demirbilek, Z. (2003). “Simulation of waves at duck (North
534 Carolina) using two numerical models.” *Coast. Eng. J.*, 45(3), 439–469.

535	List of Figures	
536	1	Sketch of the harbor model: computational domain with PML and idealization
537		of the exterior bathymetry. 27
538	2	Circular scattering with constant bathymetry: isolines of the amplification
539		factor of the numerical solution for a centered circled obstacle (left) and for
540		a not centered one (right). The dashed line indicates the position of the
541		symmetry axis. 28
542	3	Circular scattering with constant bathymetry: relative error of the amplifica-
543		tion factor ($ H - H^{\text{exact}} /H^{\text{exact}}$) for a centered circled obstacle (left) and for
544		a not centered one (right). 29
545	4	Circular scattering with constant bathymetry: convergence of the \mathcal{L}^2 norm of
546		the error, see Eq. (14), along the circle boundary. Four-order finite elements
547		are used and mean convergence rate is reported. 30
548	5	Semicircular scattering with variable bathymetry: problem statement (top)
549		and distribution of the amplification factor with an incoming wave angle $\theta =$
550		220° (bottom). The incident wave field (0.6 s of period) is also depicted
551		(top) in the PML artificial layer accordingly to each of the two cross-shore
552		bathymetries. 31
553	6	Semicircular scattering with variable bathymetry: PML influence on the am-
554		plification factor along the semicircle obstacle for a fixed finite element mesh,
555		and two different error measures: mean (15) and maximum (16) errors. In-
556		coming wave angle is $\theta = 220^\circ$ (left) and $\theta = 310^\circ$ (right). 32

557	7	Semicircular scattering with variable bathymetry: isolines for an amplification factor $H = 1$ (left column) and their corresponding isolines of correct significant digits of accuracy (right column). They are depicted in the interior area of interest defined by $D'_1 = D'_2 = 0$ for different positions of the left and right PML artificial layers. That is, each row shows the solution and the accuracy provided by a relative distance to PML of $D'_1 = 1$ (top row), $D'_1 = 2$ (middle row) and $D'_1 = 3$ (bottom row). In all cases $D'_2 = 0$	33
568	8	Mataró harbor: problem statement. Three different computational domains are depicted: largest domain (top), medium domain (bottom left) and small domain (bottom right). Sections 1-3 of interest are highlighted in the domain 1, as well as the absorbing coefficient α in Eq. (3). The solution of the amplification factor for an incoming wave direction of 225° and 6 s of period is shown in the domain 2.	34
569	9	Mataró harbor: comparison of the amplification factor on the sections of interest for short waves (6 s of period). The incoming wave directions are 270° (top row) and 225° (bottom row).	35
570	10	Mataró harbor: comparison of the amplification factor on the sections of interest for long waves (16 s of period). The incoming wave directions are 270° (top row) and 225° (bottom row).	36
571	11	Mataró harbor with the medium computational domain: the incident wave field is shown for an incoming direction of 225° for short (left) and longer waves (right).	37
572			
573			
574			
575			
576			
577			
578			

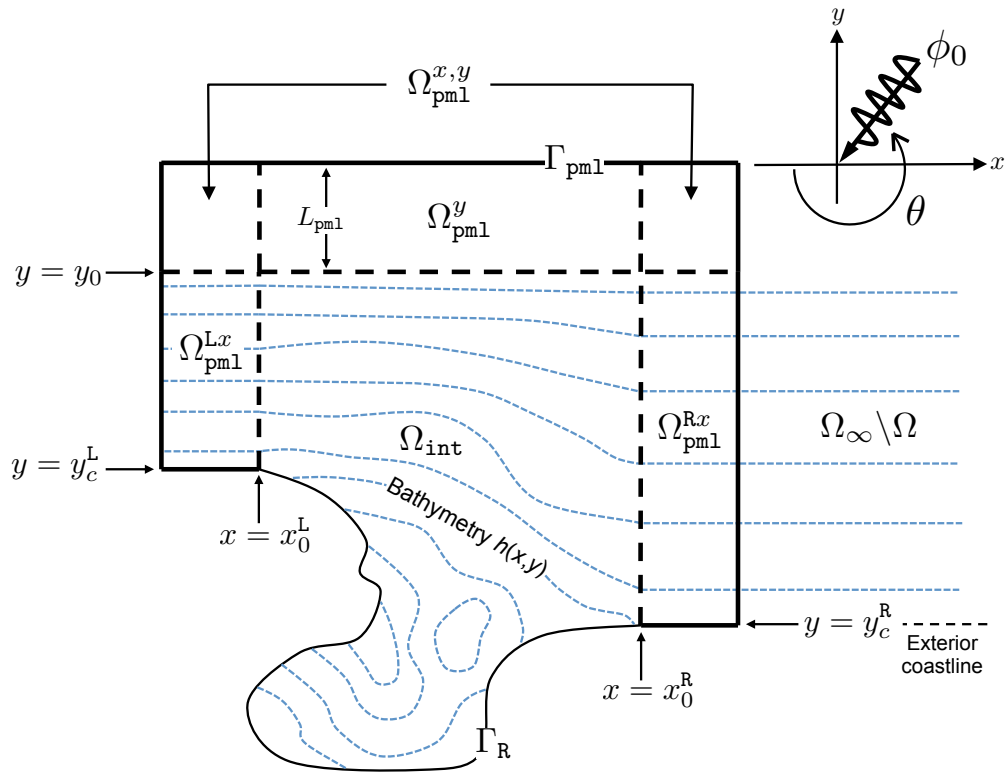


FIG. 1. Sketch of the harbor model: computational domain with PML and idealization of the exterior bathymetry.

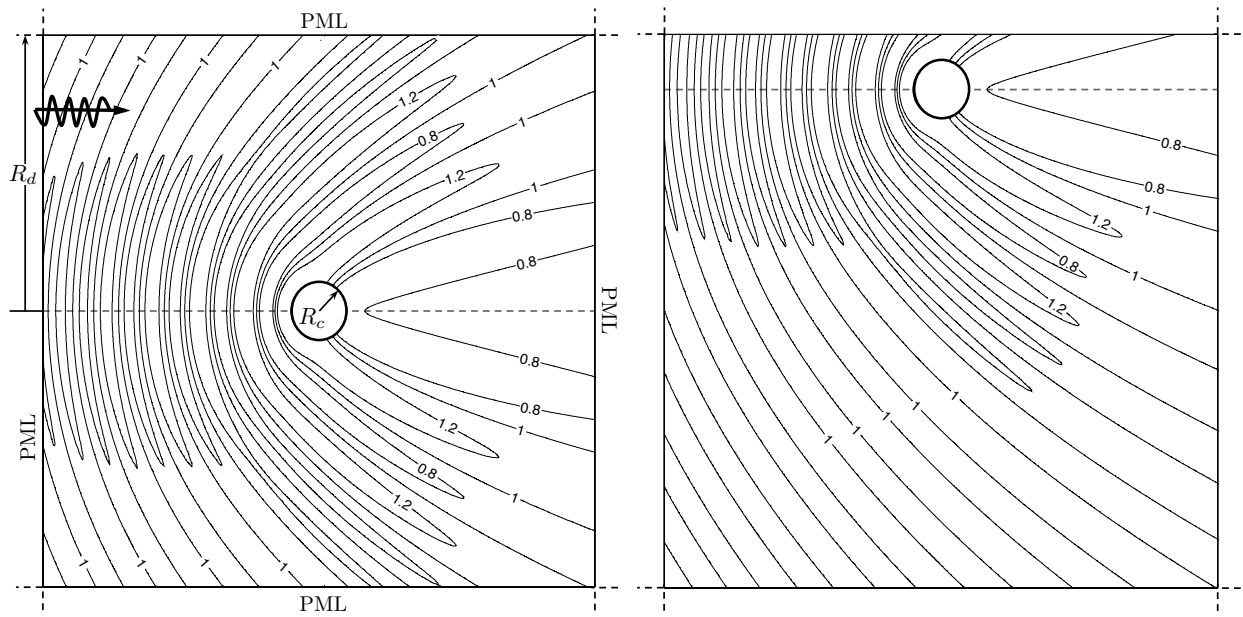


FIG. 2. Circular scattering with constant bathymetry: isolines of the amplification factor of the numerical solution for a centered circled obstacle (left) and for a not centered one (right). The dashed line indicates the position of the symmetry axis.

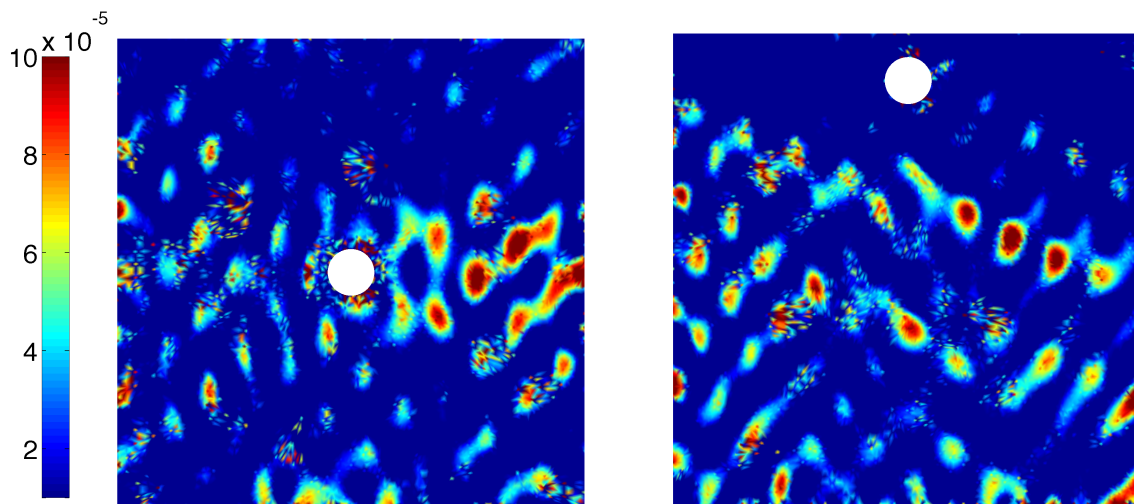


FIG. 3. Circular scattering with constant bathymetry: relative error of the amplification factor ($|H - H^{\text{exact}}|/H^{\text{exact}}$) for a centered circled obstacle (left) and for a not centered one (right).

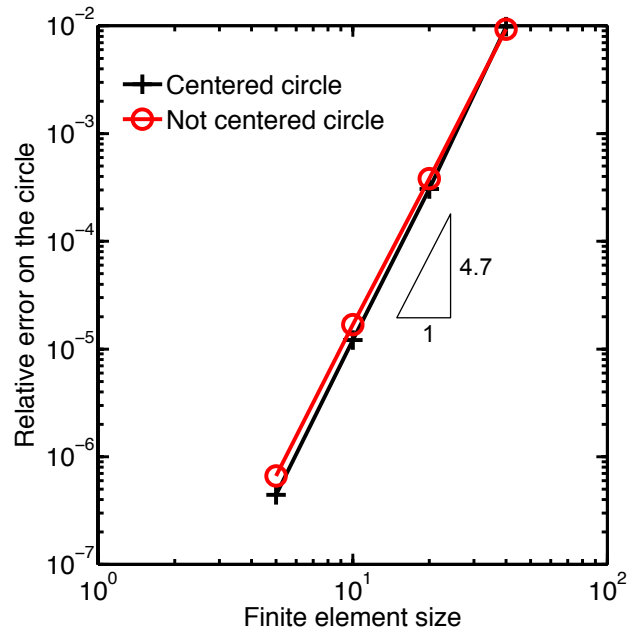


FIG. 4. Circular scattering with constant bathymetry: convergence of the L^2 norm of the error, see Eq. (14), along the circle boundary. Four-order finite elements are used and mean convergence rate is reported.

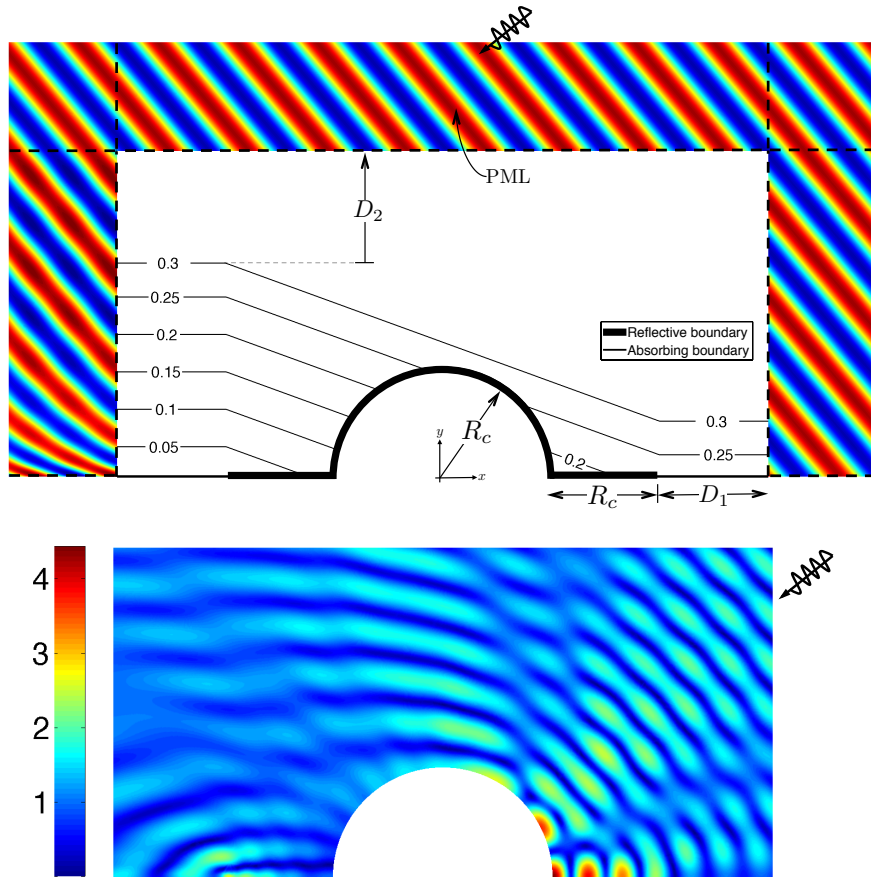


FIG. 5. Semicircular scattering with variable bathymetry: problem statement (top) and distribution of the amplification factor with an incoming wave angle $\theta = 220^\circ$ (bottom). The incident wave field (0.6 s of period) is also depicted (top) in the PML artificial layer accordingly to each of the two cross-shore bathymetries.

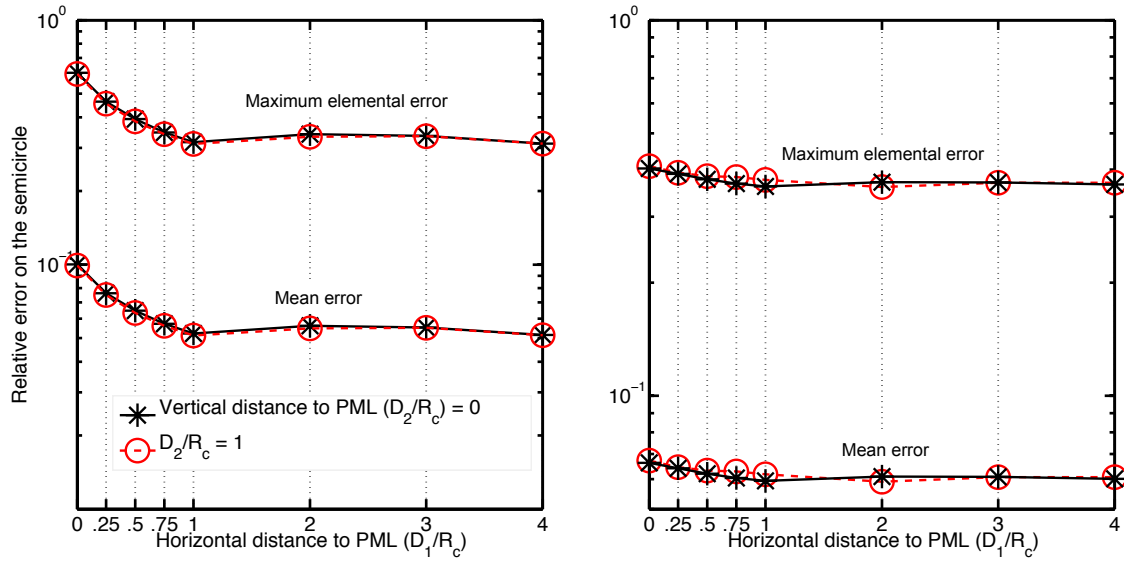


FIG. 6. Semicircular scattering with variable bathymetry: PML influence on the amplification factor along the semicircle obstacle for a fixed finite element mesh, and two different error measures: mean (15) and maximum (16) errors. Incoming wave angle is $\theta = 220^\circ$ (left) and $\theta = 310^\circ$ (right).

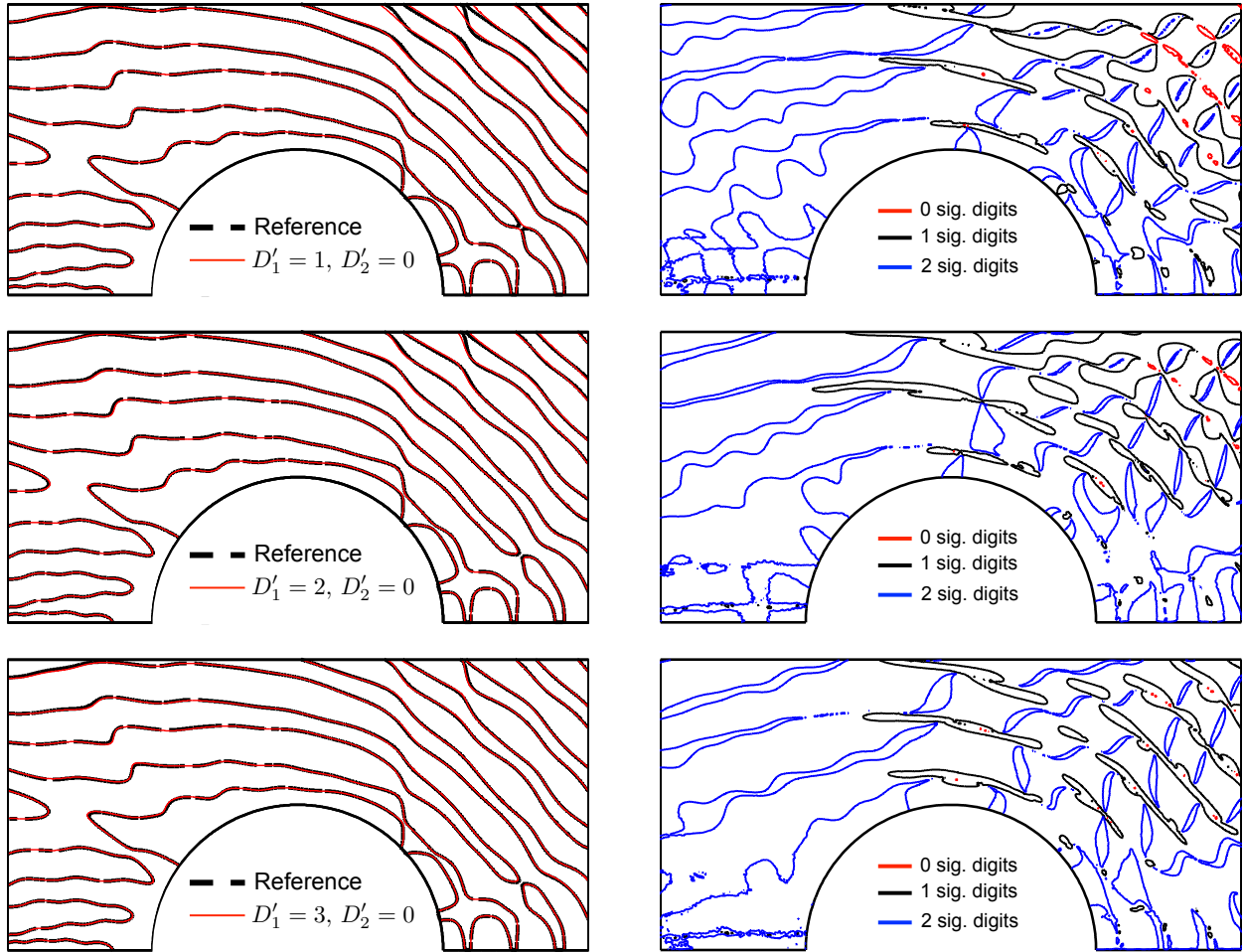


FIG. 7. Semicircular scattering with variable bathymetry: isolines for an amplification factor $H = 1$ (left column) and their corresponding isolines of correct significant digits of accuracy (right column). They are depicted in the interior area of interest defined by $D'_1 = D'_2 = 0$ for different positions of the left and right PML artificial layers. That is, each row shows the solution and the accuracy provided by a relative distance to PML of $D'_1 = 1$ (top row), $D'_1 = 2$ (middle row) and $D'_1 = 3$ (bottom row). In all cases $D'_2 = 0$.

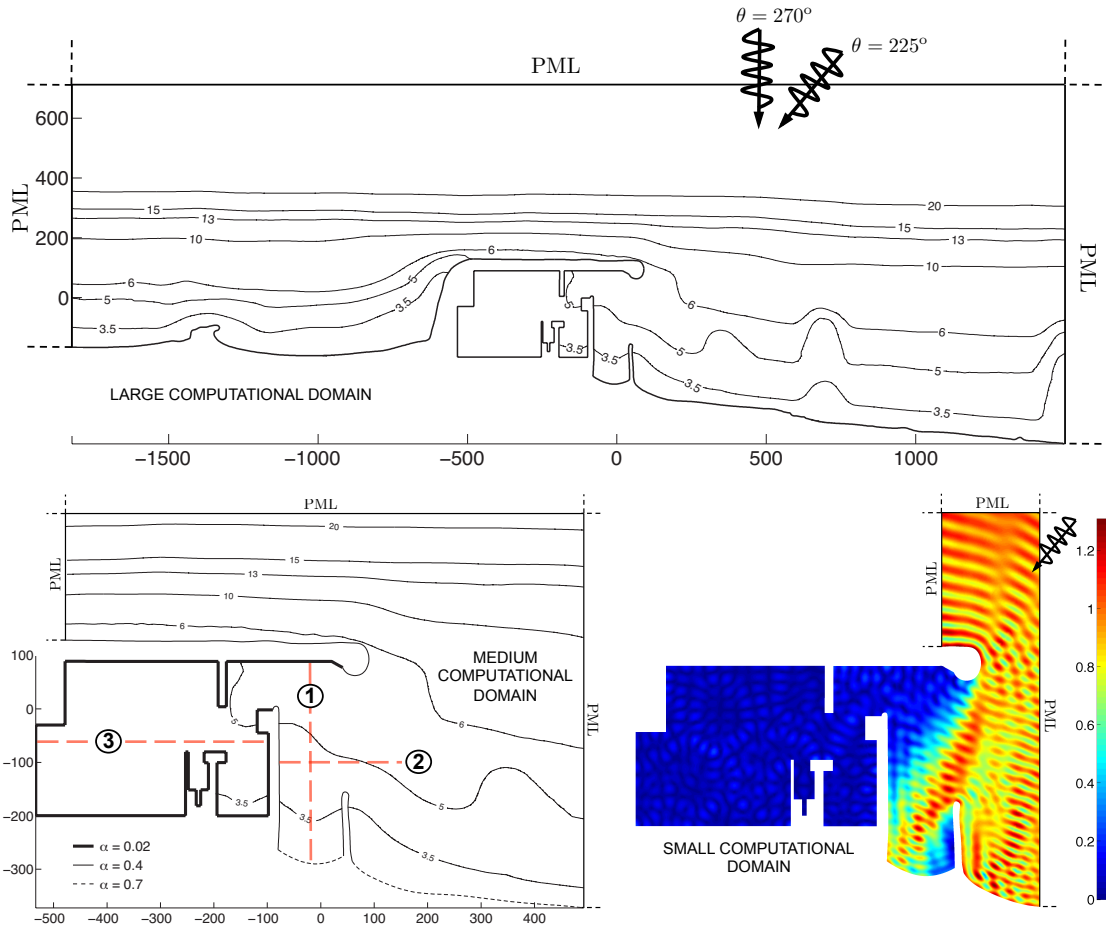


FIG. 8. Mataró harbor: problem statement. Three different computational domains are depicted: largest domain (top), medium domain (bottom left) and small domain (bottom right). Sections 1-3 of interest are highlighted in the domain 1, as well as the absorbing coefficient α in Eq. (3). The solution of the amplification factor for an incoming wave direction of 225° and 6 s of period is shown in the domain 2.

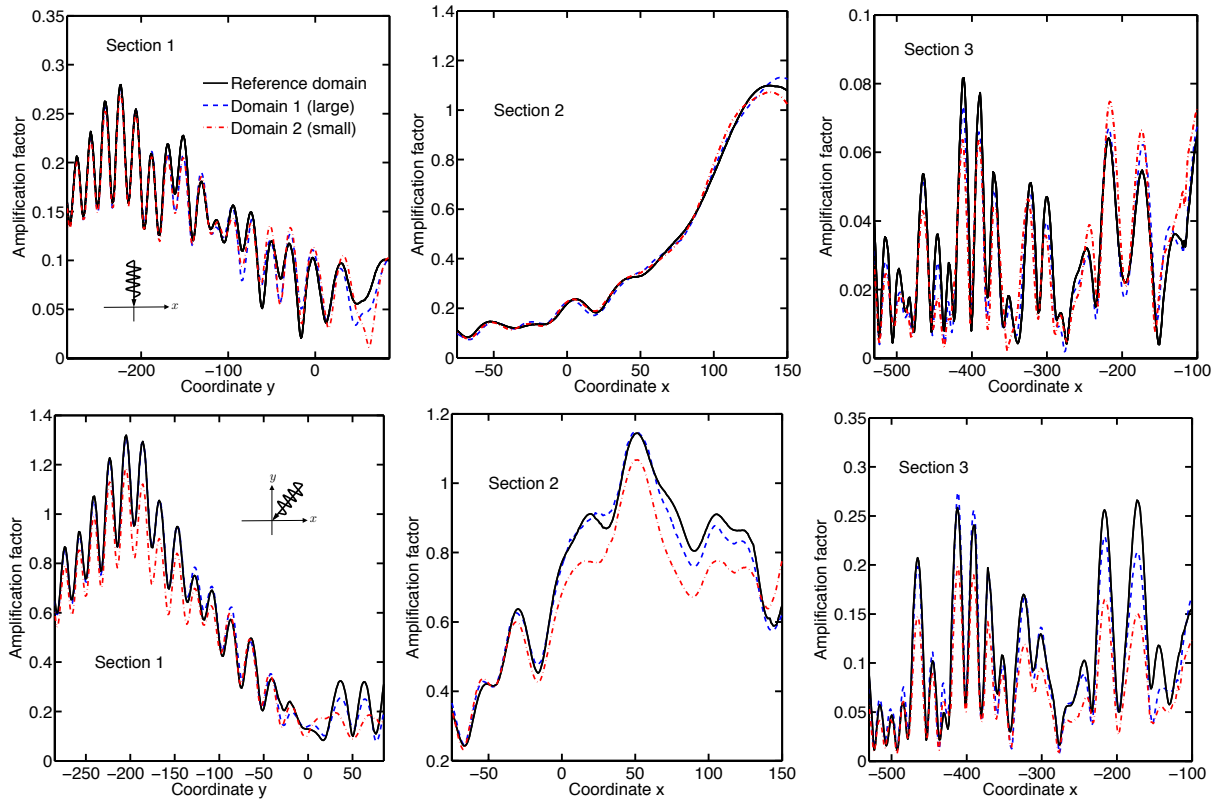


FIG. 9. Mataró harbor: comparison of the amplification factor on the sections of interest for short waves (6 s of period). The incoming wave directions are 270° (top row) and 225° (bottom row).

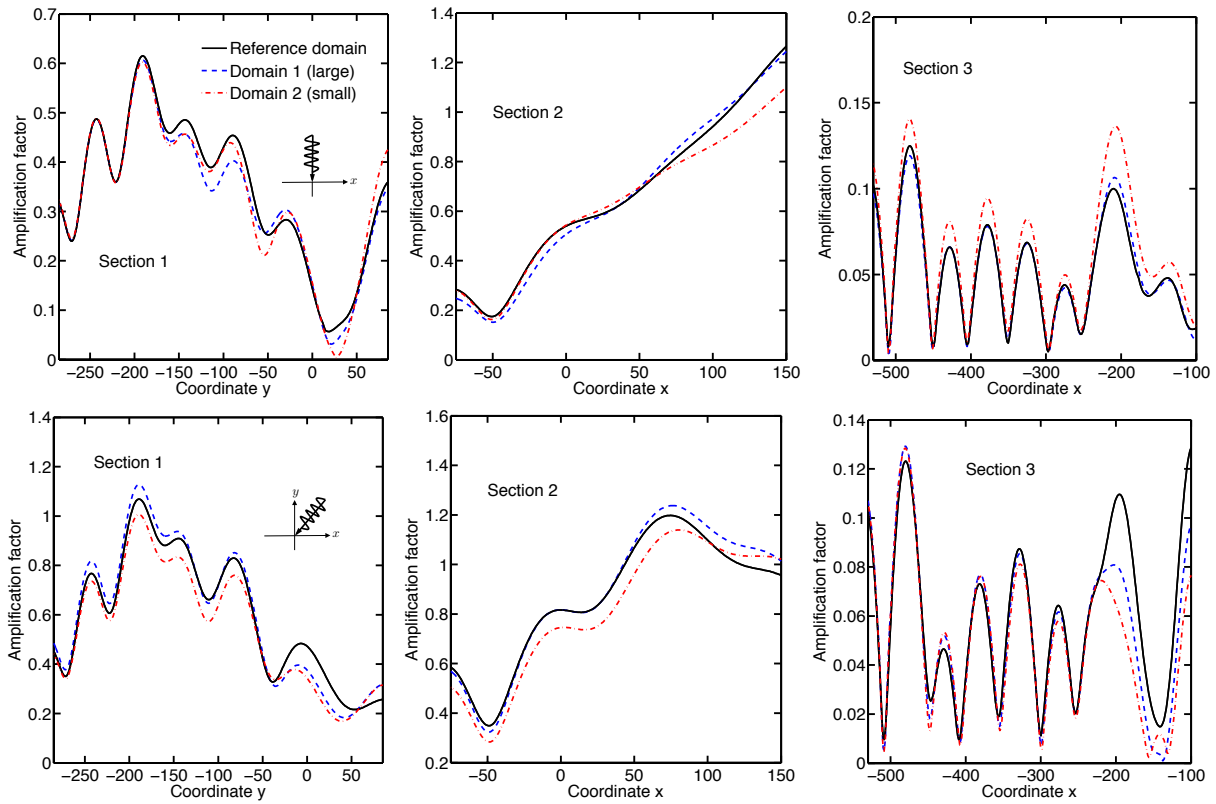


FIG. 10. Mataró harbor: comparison of the amplification factor on the sections of interest for long waves (16 s of period). The incoming wave directions are 270° (top row) and 225° (bottom row).

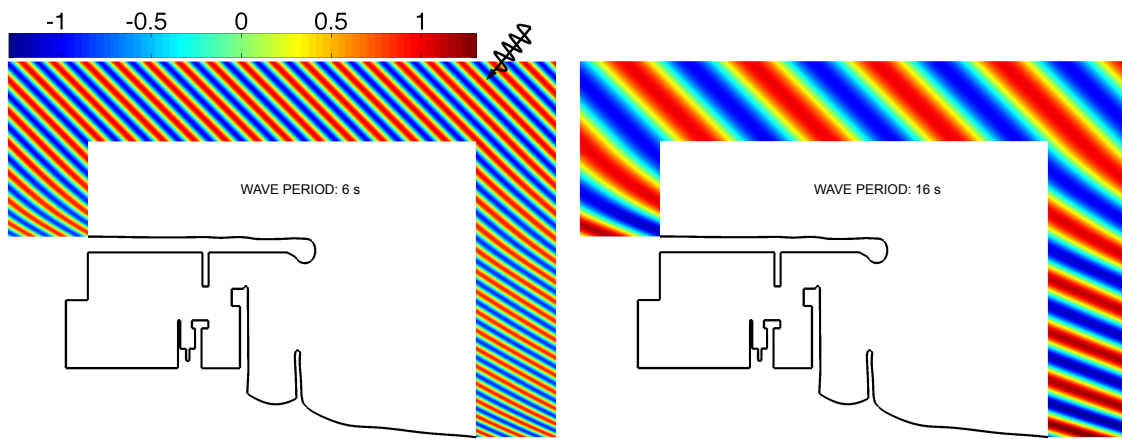


FIG. 11. Mataró harbor with the medium computational domain: the incident wave field is shown for an incoming direction of 225° for short (left) and longer waves (right).