

3 **Nitrogen inputs influence on biomass and trophic structure of ocean plankton: a study using**  
4 **biomass and stable isotope size-spectra**

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22  
23 **ABSTRACT**

24 Large scale patterns in planktonic food web structure were studied by applying continuous size-  
25 scaled models of biomass and  $\delta^{15}\text{N}$  to plankton samples, collected at 145 stations during the  
26 Malaspina-2010 Expedition across three ocean basins and including major biomes. Carbon biomass

27 and  $\delta^{15}\text{N}$  were determined in size-fractionated samples (40 to 5000  $\mu\text{m}$ ) collected by vertical hauls  
28 (0-200 m). Biomass-normalized size-spectra were constructed to summarize food web structure and  
29 spatial patterns in spectral parameters were analyzed using geographically-weighted regression  
30 analysis. Except in the northwestern Atlantic, size-spectra showed low variability, reflecting a large  
31 homogeneity in nitrogen sources and food web structure for the central oceans. Estimated predator-  
32 to-prey mass ratios  $<10^4$  and mean trophic transfer efficiency values between 16% (coastal biome)  
33 and  $>20\%$  (Trades and Westerlies biomes) suggested that oceanic plankton food webs could support  
34 a larger number of trophic levels than current estimates based on high efficiency values. The largest  
35 changes in spectral parameters and nitrogen sources were related to inputs of atmospheric nitrogen,  
36 either from diazotrophic organisms or dust deposition. These results suggest geographic  
37 homogeneity in the net transfer of nitrogen up the food web.

## 38 INTRODUCTION

39 Ocean plankton is an essential contributor to the dynamics of the earth climate system.  
40 Phytoplankton drawdown of  $\text{CO}_2$  is made at the expenses of nutrients as nitrogen, and the fate of  
41 the organic matter produced depends on the characteristics of the food web, that in turn are largely  
42 determined by the size of planktonic organisms (Legendre and Lefevre, 1995). The effects of  
43 warming and increasing stratification imply reductions in the input of nutrients from deep waters to  
44 the surface ocean and subsequent changes in primary production in large ocean areas (Behrenfeld *et*  
45 *al.*, 2006). Depending on the structure of the food web the climate effects can be amplified or  
46 modified as shown by both *in situ* (Richardson and Schoeman, 2004) and modeling studies (Chust  
47 *et al.*, 2014). Thus, determining the structure of the food web is essential for understanding the  
48 behavior of the ocean ecosystems under a changing climate.

49 Based on the dependency of most physiological processes on organism size, models relating  
50 organism abundance (or biomass) and individual size have been developed with the purpose of  
51 describing the continuous change of biomass across the whole food web (Platt and Denman, 1978;  
52 Blanco *et al.*, 1994). The so-called size-spectrum models have been used to synthesize the food web  
53 structure in comparative analysis of oceanic ecosystems (Rodriguez and Mullin, 1986; Piontkovski  
54 *et al.*, 2003; Quiñones *et al.*, 2003; San Martin *et al.*, 2006). Furthermore, they have been used to  
55 infer the maximum number of trophic levels that a particular ecosystem can support (Zhou, 2006;  
56 Basedow *et al.*, 2010). Their applicability has been questioned as their predictions generally require  
57 steady state conditions and uniformity in the size-dependent physiological rates (Poulin and Franks,  
58 2010), but some of these limitations can be overcome by models (Ward *et al.*, 2014) or the use of  
59 seasonal averages (Hunt *et al.*, 2015).

60 Stable isotope analysis provides an important tool for elucidating trophic structure. For instance,  
61 enrichment in  $^{15}\text{N}$  has been observed with increasing trophic position, thus allowing to infer trophic  
62 structure (e.g. Post, 2002). This enrichment can also be expected across plankton size classes, as  
63 pelagic food webs are strongly size-structured, the smallest organisms being generally primary  
64 producers and large organisms consumers of smaller prey (Platt and Denman, 1978). However, only  
65 a few field studies explored the implications of the variability in  $^{15}\text{N}$  with plankton size (Fry and  
66 Quiñones, 1994; Rolff, 2000; Bode *et al.*, 2007; Jennings *et al.*, 2008; Mompeán *et al.*, 2013). One  
67 of the main limitations was the difficulty in obtaining measurements over a large number of size  
68 classes representative of the different trophic levels. The adjustment of simple continuous functions  
69 was not always possible because of the frequent exceptions to the general increase of  $^{15}\text{N}$  with the  
70 average size of the organisms sampled size. For instance, plankton of total length smaller than 200  
71  $\mu\text{m}$  tend to show lower variability in average  $^{15}\text{N}$  content than plankton in larger size classes (Rolff,  
72 2000; Bode *et al.*, 2007). This is due both to the high variability at short time scales in small size  
73 classes owing to fast turnover times, thus reflecting rapid changes in the N sources (e.g. Jennings *et al.*,  
74 2008), but also to the small changes in isotopic composition observed in microbial food webs  
75 (e.g. Gutierrez-Rodriguez *et al.*, 2014). In addition, the presence of large or colonial plankton that  
76 feed on phytoplankton caused a decrease in  $^{15}\text{N}$  at large sizes, while there is a lineal increase with  
77 plankton size when these large herbivores were not present (Fry and Quiñones, 1994; Hunt *et al.*,  
78 2015). Besides, because most studies used net sieves with a logarithmic increase in mesh size, the  
79 direct comparison of  $^{15}\text{N}$  content with size did not allowed for the computation of significant linear  
80 functions in some of the samples, even when considering a relatively large range of sizes (e.g.  
81 Rolff, 2000). As for the biomass distributions, the use of logarithmic transformations and  
82 normalization of the variable of interest ( $^{15}\text{N}$  in this case) by the biomass of the organisms may lead  
83 to the construction of size-spectra that allow inferences on the trophic structure of plankton  
84 independently of the actual range of sizes measured. To our knowledge, this study represents the  
85 first application of biomass-normalized size spectra using stable isotopes.

86 Large regions of the open ocean are poorly sampled. Here, food webs are generally characterized by  
87 low nutrient concentrations, low biomass and dominance of microbial recycling of nutrients with  
88 some exceptions due to external nutrient supply (Longhurst, 2007). Studies analyzing the planktonic  
89 size structure at large oceanic scales are scarce (Piontkovstki *et al.*, 2003; Quiñones *et al.*, 2003;  
90 San Martín *et al.*, 2006) while there are no studies addressing the changes of stable nitrogen  
91 isotopes with plankton size at such scales. Meta-analyses of planktonic isotopic composition on  
92 plankton at large geographic scales have shown latitudinal patterns that were related to the sources  
93 and bioavailability of nutrients (Bowen, 2010; McMahon *et al.*, 2013). In this study we use biomass

94 and  $^{15}\text{N}$  size spectra to assess whether the type of dominant source of inorganic nutrients affects the  
95 trophic structure of plankton in the open ocean. The underlying hypothesis is that the trophic  
96 structure of plankton in the open ocean will be determined by the balance between nutrient inputs  
97 from deep or continental waters or from the atmosphere. Advective inputs can be expected to be of  
98 importance at relatively small spatial or temporal scales, as those related to mesoscale dynamics  
99 (Oschlies and Garçon, 1998), while most of the nutrients for the oligotrophic ocean are provided by  
100 transport across the pycnocline (Mouriño-Carballido *et al.*, 2011; Torres-Valdés *et al.*, 2009;  
101 Fernández-Castro *et al.* 2015). Atmospheric inputs include biological  $\text{N}_2$  fixation (Capone *et al.*,  
102 2005; Mulholland, 2007) and deposition of inorganic and organic nutrients, the latter noticeably  
103 enhanced by anthropogenic emissions (Duce *et al.*, 2008). Advection of deep nutrients is expected  
104 to lead to blooms of phytoplankton of relatively large size, high primary production and metazoan  
105 food webs, typical of most temperate and polar regions, with a strong seasonal variability  
106 (Longhurst, 2007). Most of the open ocean is expected to depend on non-seasonal, relatively small  
107 inputs of nutrients by diffusion, leading to phytoplankton of small size, low primary production and  
108 rapid remineralization of organic matter in microbial food webs. Atmospheric N fixation can be due  
109 to either large phytoplankton (as the colony-forming cyanobacteria *Trichodesmium*) or to microbial  
110 forms, but in all cases low primary production and high microbial remineralization is expected  
111 before effective transfer of the fixed N up the food web (Mulholland, 2007). Because of the  
112 dependence of N fixation on micronutrients, as Fe provided by dust deposition (Moore *et al.*, 2009),  
113 atmospheric inputs of nutrients may have also a clear seasonal component. In all cases, the trophic  
114 structure of the food web will depend on the net amount of nutrients transferred from the primary  
115 producers to the various types of consumers. This implies, by assuming a constant transfer  
116 efficiency between trophic levels (Zhou, 2006), that most of the oligotrophic open ocean would  
117 have shorter food webs than productive seasonal regions, as the biomass of pelagic predators  
118 depends on the amount of primary production (Chassot *et al.*, 2007). However, a larger number of  
119 trophic steps is expected in the oligotrophic ocean because the rapid recycling of organic matter in  
120 microbial food webs (Sommer *et al.*, 2002).

121 The objective of this study is to investigate the relationships between plankton trophic structure and  
122 the source of N at large-scale spatial scales in the ocean. For this purpose we analyzed carbon  
123 biomass and the natural abundance of stable nitrogen isotopes in a large set of size-fractionated  
124 plankton samples from the research expedition Malaspina-2010 across three ocean basins.

## 125 **METHODS**

### 126 **Plankton sampling and analysis**

127 Plankton samples were collected during Malaspina-2010 expedition between December 2010 and  
128 July 2011 (Fig. 1). The expedition employed two oceanographic ships to make observations and  
129 collect water, seston and plankton samples across three major ocean basins  
130 (<http://metamalaspina.imedea.uib-csic.es/geonetwork/srv/en/main.home>). In this study plankton  
131 samples from 145 stations were considered. Samples were collected by vertical hauls of bongo-type  
132 nets (30 cm diameter, 40  $\mu\text{m}$  mesh size and 50 cm diameter, and 200  $\mu\text{m}$  mesh size) between 200 m  
133 depth and the surface during early morning hours. Plankton was size-fractionated using sieves of  
134 200, 500, 1000, 2000 and 5000  $\mu\text{m}$ , collected on pre-weighted glass-fiber filters and oven dried  
135 (60°C, 24 h) on board. Large gelatinous organisms were removed and analyzed separately (Molina-  
136 Ramirez et al., 2015). In addition, sample aliquots were preserved in formalin (4% final  
137 concentration) for later determination of the abundance of trichomes of the N-fixing  
138 cyanobacterium *Trichodesmium*. Counts were made using a semiautomatic image analysis flow-  
139 through system (FlowCam).

140 Biomass was later determined in the laboratory for each size fraction as carbon (C) and nitrogen  
141 content using an elemental analyzer (Carlo Erba CHNSO 1108). No acidification treatment to  
142 remove carbonates was applied prior to analysis of carbon, but molar C:N values (mean  $\pm$  se = 4.8  $\pm$   
143 1.39, n = 145) were typical of planktonic organic matter (Søreide et al., 2007) and indicated that the  
144 possible overestimation of C due to inorganic carbon would be small. Natural abundance of stable  
145 nitrogen isotopes was determined using a mass spectrometer (Finnigan Mat Delta Plus) coupled to  
146 the elemental analyzer. Nitrogen stable isotope abundance were expressed as  $\delta^{15}\text{N}$  (‰) relative to  
147 atmospheric nitrogen (Coplen, 2011). International Atomic Energy Agency (IAEA) USGS40 and L-  
148 alanine isotope standards were analyzed with the samples, along with internal acetanilide and  
149 sample standards. Precision ( $\pm$ se) of replicate determinations of standards and samples was <0.1‰  
150 (n= 2 to 6) and <0.3‰ (n= 5 to 10), respectively. The analytical offset between certified and  
151 measured values was <0.1‰. All isotopic determinations were made in the Servicio de Análisis  
152 Instrumental of the Universidade da Coruña (Spain). Further details on the sampling and analysis  
153 can be found in Moreno-Ostos (Moreno-Ostos, 2012) and Mompeán *et al.* (Mompeán *et al.*, 2013).

#### 154 **Oceanographic conditions *in situ***

155 Hydrographic information was obtained from CTD-rosette casts at the same stations and  
156 chlorophyll-a (Chl<sub>a</sub>) was determined from acetonic extracts of phytoplankton collected at up to 8  
157 discrete depths in the photic layer (>0.1% of surface photosynthetically active irradiance). Here  
158 Chl<sub>a</sub> values were integrated in the photic layer as a proxy for phytoplankton biomass. In addition,  
159 the vertical extent of phytoplankton was estimated by the depth of the chlorophyll maximum

160 (DCM). As nutrient concentrations and fluxes were not available for all the sampled stations  
161 (Fernández-Castro et al., 2015), several proxies for nutrient supply from deep waters were  
162 employed. First, the DCM itself, as it was negatively correlated with the flux of nitrate (Fig. 1S).  
163 Second, the stratification of the upper water column (Behrenfeld et al., 2006). estimated by the  
164 mixing layer depth (MLD), computed using a density difference criterion ( $\Delta\delta = 0.125 \text{ kg m}^{-3}$  with  
165 respect to surface values), and the mean squared Brunt-Väisälä frequency ( $N^2$ ) values computed at 1  
166 m intervals for the upper 200 m. Additional details on the sampling and on the analytical methods  
167 employed can be found in Moreno-Ostos (Moreno-Ostos, 2012).

## 168 **Satellite observations**

169 As the observed plankton properties (i.e. biomass, stable isotopes) at each station would be the  
170 consequence not only of local conditions but also of general oceanographic conditions prevailing  
171 over a certain amount of space and time, additional variables, estimated from satellite observations,  
172 were also considered. Annual averages of primary production for 2010 (PP,  $\text{mg C m}^{-2} \text{ d}^{-1}$ ) were  
173 generated by averaging monthly data for primary production downloaded from the Ocean  
174 Productivity website (<http://www.science.oregonstate.edu/ocean.productivity/index.php>) from the  
175 grid ( $0.17^\circ \times 0.17^\circ$ ) closest to each station position.

176 Dust deposition, as a proxy for atmospheric inputs of key nutrients for primary production (e.g. Fe  
177 or P), was also estimated from Aqua-MODIS Aerosol Optical Depth at 550 nm and Aerosol Small  
178 Mode Fraction data provided by the Giovanni online data system (NASA Goddard Earth Sciences).  
179 These data were retrieved from a grid of  $1^\circ$  resolution and centered at the closest location to each  
180 station and combined with wind speed derived from AVISO  
181 (<http://las.aviso.oceanobs.com/las/getUI.do>) to estimate the monthly average atmospheric dust  
182 column concentration (MDU,  $\text{g m}^{-2}$ ) at each station (Kaufman *et al.*, 2005).

## 183 **Creation of size spectra**

184 Biomass and  $\delta^{15}\text{N}$  values by size fractions of plankton were combined in the form of linear  
185 regressions with the median size of individuals in the corresponding size fraction. This median size  
186 was estimated as the geometric mean of the nominal size (length) of the sieves employed for each  
187 class, and further calculation of carbon biomass by using conversion factors (Rodriguez and Mullin,  
188 1986). Due to the large variability in the values, a single regression including all stations sampled  
189 was not possible (Fig. 1S). Logarithmic transformations ( $\log_2$ ) and normalization by the width of  
190 the interval of individual biomass were applied to obtain functions independent from size-class  
191 width (Blanco *et al.*, 1994, Zhou, 2006). The resulting biomass-normalized lines (obtained by least

192 squares regression) characterized the size distribution of biomass or  $\delta^{15}\text{N}$  by two parameters: the  
193 intercept ( $C_a$ ,  $\delta^{15}\text{N}_a$ ), a proxy for the values in the first size class, and the slope ( $C_b$ ,  $\delta^{15}\text{N}_b$ ),  
194 indicative of the rate of change in the values with increasing individual size (Fig. 2). Individual  
195 values of spectral regression parameters for each station are provided in the Supplement (Table 1S).  
196 The normalization procedure can be reversed to obtain un-normalized values of the slope that can  
197 be further used to develop several descriptors of food web properties (Blanco et al., 1994). For  
198 instance, we computed the number of trophic levels (NTL) from  $C_b$  (Zhou, 2006), the predator-to-  
199 prey mass ratio (PPMR) from the slope  $m$  of the un-normalized  $\delta^{15}\text{N}$  spectrum (Jennings *et al.*,  
200 2001), and the trophic transfer efficiency (TTE) from PPMR (Barnes *et al.*, 2010). Where:

$$201 \quad \text{NTL} = - (1+0.7) / (C_b 0.7)$$

202 by assuming a trophic efficiency (i.e. the ratio of production of predator to its prey) of 0.7 (Zhou,  
203 2006),

$$204 \quad \text{PPMR} = 2^{2.2/m}$$

205 by assuming a  $\delta^{15}\text{N}$  enrichment between trophic levels of 2.2‰, appropriate for plankton,  
206 ammonotelic organisms which show lower  $\delta^{15}\text{N}$  enrichment than consumers excreting urea  
207 (Vanderklift and Ponsard, 2003; Hunt et al., 2015), and:

$$208 \quad \text{TTE} = \text{PPMR}^{-1.05+0.75}$$

209 by assuming a time averaged un-normalized biomass slope of -1.05 (Barnes et al., 2010; Hunt et al.,  
210 2015).

211 In addition, the trophic position of the largest plankton size class ( $\text{TP}_{2000}$ ) was computed from their  
212  $\delta^{15}\text{N}$  values ( $\delta^{15}\text{N}_{2000}$ ) as:

$$213 \quad \text{TP}_{2000} = \text{TP}_{40} + \frac{\delta^{15}\text{N}_{2000} - \delta^{15}\text{N}_{40}}{2.2}$$

214 where  $\text{TP}_{40}$  is the trophic position of the 40-200  $\mu\text{m}$  size class used as the reference baseline by  
215 assuming that it contained a mixture of phytoplankton and microzooplankton ( $\text{TP}_{40} = 1.5$ ), and a  
216 constant  $\delta^{15}\text{N}$  enrichment between adjacent trophic levels of 2.2‰ (Hunt et al., 2015).

## 217 **Spatial analysis**

218 Stations were first grouped by biomes (Longhurst, 2007) (Fig. 1) to investigate large scale  
219 differences in biomass and spectral parameters using ANOVA and Dunnett-C post-hoc tests.

220 Visualization and further analysis of trends at various spatial scales were made using the tools  
221 provided by the package Spatial Analysis in Macroecology (SAM V 4.0) which allowed mapping  
222 and computation of various spatial statistics from metrics descriptive of spatial autocorrelation to  
223 advanced spatial regression (Rangel *et al.*, 2010). Spatial autocorrelation of variables was analyzed  
224 by means of correlograms of the Moran's I coefficient computed for groups of samples (stations) of  
225 increasing spatial distance. The relationships between spectral parameters and environmental  
226 variables were investigated using principal components analysis (PCA) to determine the main  
227 correlations. Then, we selected the most representative environmental variables to construct  
228 regression models explanatory of each of the spectral parameters at each station using  
229 Geographically Weighted Regression (GWR). This method, specifically designed to deal with  
230 spatial autocorrelation and lack of stationarity (Rangel *et al.*, 2010), was used to compute a series of  
231 local regressions, one for each station location, between the independent variable (i.e. spectral  
232 parameters) and the explanatory variables (i.e. selected environmental variables) taking into account  
233 the information from the surrounding stations weighted by a spatial function. In this case we used a  
234 moving spatial window with an adaptive Kernel of 15% of neighboring stations optimized using the  
235 Akaike's Information Criterion. The explanatory power of GWR ( $r^2$ ) for each spectral parameter  
236 was in general higher than those of ordinary least squares, as measured by ANOVA tests (Rangel *et al.*,  
237 *et al.*, 2011).

## 238 **RESULTS**

### 239 **Spatial distributions**

240 Biomass and  $\delta^{15}\text{N}$  showed an uneven distribution across the ocean as illustrated by the maps for the  
241 40-200  $\mu\text{m}$  size-class (Fig. 3). In contrast, the values for the parameters characterizing biomass and  
242  $\delta^{15}\text{N}$  spectra showed in general low variability, resulting in a relatively homogeneous distribution at  
243 large spatial scales (Fig. 4, Fig. 3S). An exception was the Subtropical North Atlantic region (cruise  
244 legs 7 and 8), mostly in the Westerlies biome. Here, the intercepts for both biomass and  $\delta^{15}\text{N}$   
245 spectra ( $C_a$  and  $\delta^{15}\text{N}_a$ ) reached in general lower values than in other regions, while the slopes ( $C_b$   
246 and  $\delta^{15}\text{N}_b$ ) showed high variability. This latter pattern is a response to the low  $^{15}\text{N}$  values measured  
247 in the 40-200  $\mu\text{m}$  size-class (Fig. 3) in the subtropical regions thus producing an overall negative  
248 correlation with latitude ( $r=-0.507$ ,  $P<0.001$ ,  $n=145$ ). The geographic and temporal distribution of  
249 samples also affected in complex ways to these distributions, as there were positive correlations  
250 between  $\delta^{15}\text{N}_b$  and either latitude ( $r=0.331$ ,  $P<0.001$ ) or sampling date ( $r=0.202$ ,  $P<0.05$ ). When  
251 grouped by biomes, the differences were significant in the Westerlies (including the southern  
252 Australia region) for all parameters and variables, except for  $\delta^{15}\text{N}_b$  (Table 1). Thus, the lower



253 biomass and  $\delta^{15}\text{N}$  values at the base of the planktonic food web in the Westerlies biome ( $C_{40-200}$  and  
254  $\delta^{15}\text{N}_{40-200}$ ) affected the food web structure represented by size spectra.

255 The trophic position of the largest plankton size-class sampled was significantly correlated with the  
256 slope of the  $\delta^{15}\text{N}$  spectra (Fig. 4S) and reached values generally below 2, particularly for the  
257 samples in the coastal biome (Table 2). In contrast, estimates using the biomass spectra resulted in a  
258 maximum number of ca. 3 trophic levels. Mean values of PPMR were below 1000 and no  
259 significant differences were found between biomes, while mean TTE was 16% for the coastal  
260 biome and >20% for the Trades and Westerlies biomes (Table 2).

261 The distribution of annual primary production (Fig. 3S f) was also quite homogeneous across the  
262 sampled ocean regions, while variables indicative of different nutrient inputs, as the abundance of  
263 *Trichodesmium*, the depth of the chlorophyll maximum or the atmospheric dust deposition were  
264 more heterogeneously distributed (Fig. 5, Fig. 3S). *Trichodesmium* abundance was present in all  
265 regions but showed in general high values in the North Atlantic (legs 7 and 8). The chlorophyll  
266 maximum was deepest in the central regions of all ocean basins. The distribution of other variables  
267 as the depth of the mixing layer or the Brunt-Väisälä frequency (not shown) had also variability at  
268 intermediate and large spatial scales but otherwise they were significantly correlated with the depth  
269 of the chlorophyll maximum ( $r = -0.355$ ,  $P < 0.01$ ,  $n = 137$ ). Relatively high dust deposition was  
270 estimated in the North Pacific and in the North Atlantic, but in the latter there were large differences  
271 between the values corresponding to different cruise legs in this region. For instance, the largest  
272 dust deposition corresponded to the spring-summer leg and the lowest to the winter leg (Fig. 5c,  
273 Fig. 3S g).

#### 274 **Correlations of spectral and other variables**

275 We found correlation between the intercepts of each spectrum type (biomass or  $\delta^{15}\text{N}$ ) and the  
276 values measured at the smallest size class (40-200  $\mu\text{m}$ ), as expected. There was also a negative  
277 correlation between the slopes of biomass and  $\delta^{15}\text{N}$  spectra. All these correlations can be  
278 characterized by the angle between vectors in the space of the two main components of the PCA,  
279 explaining 46.3% of total variance (Fig. 6). It should be highlighted the negative correlations  
280 between  $\delta^{15}\text{N}_a$  and the depth of the chlorophyll maximum (DCM) and the abundance of  
281 *Trichodesmium*, and the positive correlations between  $C_a$  and chlorophyll and primary production.  
282 In turn,  $\delta^{15}\text{N}_b$  showed positive correlation with mean dust deposition and DCM while  $C_b$  was  
283 negatively correlated with DCM. Interestingly, DCM can be considered a better index of nutrient  
284 inputs across the thermocline than the depth of the mixing layer (MLD), as it was positively

285 correlated with MLD and also with chlorophyll-*a*, primary production, biomass and  $\delta^{15}\text{N}$  of the 40-  
286 200  $\mu\text{m}$  size-class.

### 287 **Geographically weighted regression models**

288 From the global correlation and PCA results we selected the independent variables DCM (as an  
289 index of the vertical transport of nutrients across the thermocline), *Trichodesmium* abundance (as a  
290 proxy for atmospheric nitrogen fixation) and MDU (as a proxy for other atmospheric nitrogen  
291 inputs) to construct regression models explanatory of the spectral parameters ( $C_a$ ,  $C_b$ ,  $\delta^{15}\text{N}_a$ ,  $\delta^{15}\text{N}_b$ )  
292 at local scale using GWR models (Table 3). Most GWR models were significant ( $P < 0.001$ ) and  
293 only the GWR model for  $C_b$  resulted non significant ( $P = 0.070$ ) and with less explanatory power  
294 than ordinary least squares (the latter with a relative decrease of 33.035 in the Akaike Information  
295 Criterion). The spatial autocorrelation of spectral variables and residuals indicated by Moran's I  
296 values (Fig. 7) showed that the spatial structure was better represented at all spatial scales in the  
297 case of  $\delta^{15}\text{N}$  models, with lower residuals than models for biomass. The median values of the model  
298 parameters indicated that the distribution of the dependent variables was determined mainly by a  
299 constant value modulated by the coefficients of the independent variables (Table 3). *Trichodesmium*  
300 abundance and DCM had marginal influence in general (median values close to 0) but their range of  
301 values indicated some local influence. In contrast, median values for MDU coefficients ranked  
302 second after the constant values, suggesting that this variable was important for determining the  
303 shape of  $\delta^{15}\text{N}$  spectra and the biomass at the base of the food web ( $C_a$ ).

### 304 **DISCUSSION**

305 The results of this study showed that the size spectra of stable nitrogen isotopes added a new  
306 dimension to characterize the structure of planktonic food webs across the ocean. The parameters  
307 defining the size spectrum at each particular region were correlated with the main sources of  
308 nutrients for primary production, as the vertical transport across the thermocline and atmospheric  
309 inputs. Thus, the use of stable isotope spectra complement the information provided by biomass  
310 spectra, which were already established as an useful tool for describing and modeling the structure  
311 and function of pelagic ecosystems (Rodriguez and Mullin, 1986; Quiñones *et al.*, 2003; San Martin  
312 *et al.*, 2006; Zhou, 2006). The large number of observations collected across three ocean basins in  
313 our study suggests that the resulting relationships are of general application to all pelagic  
314 environments.

### 315 **Isotope size-spectra in the central region of the ocean**

316 The samples used in this study do not cover all possible oceanic conditions as most of the  
317 observations were made in central areas of the ocean. Also, temporal variability (e.g. seasonality)  
318 was not taken into account for most regions. However, three of the four ocean biomes were covered  
319 (only the polar biome was not included in the study), and the range of primary production values  
320 was representative of most of the open ocean, even including some data from the most productive  
321 upwelling regions (e.g. Benguela). In addition, the observations across the subtropical North  
322 Atlantic were made in two different seasons (winter-early spring and late spring-summer) thus  
323 providing some hints for temporal variability, at least in this region. This is supported by the overall  
324 correlations with the sampling date for either  $\delta^{15}\text{N}_{40}$  or  $\delta^{15}\text{N}_b$ . In agreement with the low relative  
325 importance of seasonality for plankton in the tropical and subtropical ocean (Longhurst, 2007), we  
326 found small differences in size-spectra in this region, even when there were large differences in dust  
327 inputs between transects made in different seasons (Fig. 5 c, Fig. 3S g). Short term variability (e.g.  
328 those caused by nycthemeral migrations of plankton between deep and surface layers) is likely to  
329 have a small effect on our results as the sampling was conducted approximately at the same period  
330 of the day through all cruises.

331 The range of individual sizes considered could also influence the estimation of the parameters of the  
332 spectra (Blanco *et al.*, 1994). Besides, a large fraction of the primary production in the central  
333 regions of the ocean is due to phytoplankton cells much smaller than the sizes considered in our  
334 samples (Marañón, 2009). Although we employed only 5 discrete size-classes to construct the  
335 spectra, they included organisms ranging from 10 ng C to 0.5 mg C individual weight (i.e. four  
336 orders of magnitude). An inspection of some of the samples collected indicated a good  
337 representation of phytoplankton (including diatoms, dinoflagellates, phytoflagellates and  
338 cyanobacteria), and zooplankton filter-feeders and predators (Mompeán *et al.*, 2013). This wide  
339 range of organisms, along with the robust character of biomass-normalized size-spectra that allows  
340 for estimations of properties of the plankton community outside the actual range of fitted sizes  
341 (Blanco *et al.*, 1994), supports the validity of our estimations for the plankton food web. In the case  
342 of  $\delta^{15}\text{N}$ , previous studies showed that the spectra are continuous across the whole pelagic food web  
343 (Jennings *et al.*, 2001; Bode *et al.*, 2007; Barnes *et al.*, 2010; Hunt *et al.*, 2015).

#### 344 **Size-spectra and nutrient sources**

345 When considering biome scales, our results show large homogeneity in the parameters defining the  
346 size spectra. Only the Westerlies biomes resulted with lower values of intercept and slope in the  
347 biomass spectra, and also lower intercept values of the  $\delta^{15}\text{N}$  spectra, than those computed for the  
348 Trades and Coastal biomes. Such homogeneity in the size structure of the open ocean plankton can

349 be interpreted in terms of the main controls of the productivity. For most ocean regions, primary  
350 production is controlled by the availability of nutrients in the photic layer, and the main inputs  
351 depend on eddy diffusion through the pycnocline (Fernández-Castro et al., 2015) or advection of  
352 subsurface waters (Longhurst, 2007). These are likely the main fertilization mechanisms for the  
353 Trades and Coastal biomes, including productive upwelling in equatorial and coastal waters, while  
354 most of the subtropical oceans are occupied by oligotrophic gyres with reduced nutrient inputs and  
355 productivity (Behrenfeld *et al.*, 2006). Nevertheless, various mechanisms can provide local nutrient  
356 inputs, as mesoscale turbulence (Oschlies and Garçon, 1998), lateral transport from productive  
357 regions (Torres-Valdés *et al.*, 2009), atmospheric deposition (Duce *et al.*, 2008) and biological  
358 fixation of atmospheric nitrogen (Capone *et al.*, 2005). Notably, the subtropical North Atlantic is a  
359 region within the Westerlies biome where the fixation of nitrogen by diazotrophic plankton is  
360 favored by the deposition of Fe and phosphorus in dust particles (Capone *et al.*, 2005; Moore *et al.*,  
361 2009; Fernández *et al.*, 2010). Notwithstanding the relative small fixation rates recorded in tropical  
362 and subtropical waters of the South Atlantic, they also contributed to a large fraction of total  
363 nitrogen inputs when compared to those from eddy diffusion (Mouriño-Carballido *et al.*, 2011).  
364 Only in the vicinity of upwelling areas the contribution of deep water nitrogen produced a traceable  
365  $\delta^{15}\text{N}$  signal in plankton (e.g. Fernández et al., 2010; Hauss et al., 2013; Mompeán et al., 2013) as  
366 observed in the increase of  $\delta^{15}\text{N}_{40}$  from east to west in legs 7 and 8 (Fig. 3).

367 Our results show how the importance of these fertilization mechanisms translates into a significant  
368 correlation between indices of different nutrient sources and plankton biomass structure, the latter  
369 represented in this study by the parameters of the size spectra. Nutrient inputs from subsurface  
370 waters are indicated by the DCM, the layer where phytoplankton biomass accumulates because  
371 growth is maximized by a compromise between light levels (from above) and nutrient inputs (from  
372 below). The DCM, generally deeper when vertical advection/diffusion is low (e.g. at the centre of  
373 oligotrophic gyres) and shallower when is high (e.g. near upwelling areas), also could be seasonally  
374 affected. Previous studies have revealed that DCM is a better proxy for productivity than  
375 instantaneous nutrient concentrations in the surface layer (Mouriño *et al.*, 2004), and estimations of  
376 the vertical flux of nitrate during the Malaspina-2010 expedition (Fernández-Castro et al., 2015)  
377 were significantly correlated with the DCM (Fig. 1S). In the case of atmospheric inputs, biological  
378 nitrogen fixation is generally associated to the abundance of the colony-forming *Trichodesmium*,  
379 although unicellular diazotrophs may be also important, as found in most of the studied stations  
380 across the North Atlantic (Benavides *et al.*, 2016). Dust deposition favors nitrogen fixation, as it  
381 provides the necessary additional micronutrients (Moore *et al.*, 2009; Fernández *et al.*, 2010), but it  
382 also introduces significant nitrogen and phosphorus amounts (Morin *et al.*, 2009). Thus, the main

383 nutrient inputs can be used to model the biomass at the base of the food web, represented in our  
384 model by  $C_a$ .

### 385 **Size-spectra and trophic structure**

386 To our knowledge, this study represents the first application of biomass-normalized size spectra  
387 using stable isotopes. The increase in  $\delta^{15}\text{N}$  with organism size is a general rule in pelagic food webs  
388 (Jennings *et al.*, 2001; Bode *et al.* 2007; Jennings *et al.* 2008; Barnes *et al.*, 2010; Hunt *et al.*,  
389 2015). However previous studies of  $\delta^{15}\text{N}$  in different size classes of plankton failed to demonstrate  
390 a regular increase in  $\delta^{15}\text{N}$  with size, as there were many exceptions both at large and small organism  
391 sizes (Fry and Quiñones, 1994; Rolff, 2000, Bode *et al.*, 2007; Landrum *et al.*, 2011, Mompeán *et*  
392 *al.*, 2013). The planktonic exception can be attributed to the presence of large herbivores (as the  
393 colony-forming salps) but also to the different turnover times of biomass in organisms of different  
394 size (Jennings *et al.*, 2008). Besides, there are evidences of low or null enrichment in  $\delta^{15}\text{N}$  within  
395 the microbial food web (Gutierrez-Rodriguez *et al.*, 2014). The normalization procedure applied in  
396 this study overcomes the lack of a regular increase in each local spectrum computed with raw  
397 values, as in previous studies. The biomass-normalized spectrum allows for a synthetic  
398 representation of the continuous distribution of the variable of interest across organism sizes by  
399 using only two parameters (Zhou, 2006). Besides, the reversal in the normalization (Blanco *et al.*,  
400 1994) facilitates the determination of the values of the un-normalized slopes employed in  
401 estimations of PPMR and TTE (Barnes *et al.*, 2010). Such un-normalized slopes cannot be obtained  
402 in a larger number of cases using raw values (e.g. Fig. 2S). In this study, changes in the intercept of  
403 the  $\delta^{15}\text{N}$  spectrum were related to the dominant source of nitrogen for primary producers, while  
404 changes in the slope were indicative of the overall isotopic enrichment along the food web.

405 Low intercepts are thus expected in areas with significant atmospheric inputs, either by diazotrophy  
406 or by inputs of anthropogenic nitrogen, as these sources have  $\delta^{15}\text{N}$  values near zero (Morin *et al.*,  
407 2009). This was the case of our observations in the Westerlies biome that had lower  $\delta^{15}\text{N}_a$  than the  
408 Trades and Coastal biomes (Table 1), in agreement with the large importance of atmospheric  
409 nitrogen fixation in the former. However, we did not find significant differences in  $\delta^{15}\text{N}_b$  between  
410 biomes and the detailed distribution of this parameter was very homogeneous across the ocean (Fig.  
411 4d). This suggests that the transfer of nitrogen up the plankton food web proceeds, in general, at  
412 similar rates in all ocean regions. The exception was found in the two legs crossing the subtropical  
413 N Atlantic, where a large variability in  $\delta^{15}\text{N}$  spectra was found in coincidence with relatively high  
414 amounts of *Trichodesmium* and other diazotrophs (Benavides *et al.*, 2016).

415 The relative homogeneity in the structure of oceanic plankton food webs suggested by  $\delta^{15}\text{N}_b$  is  
416 supported by the low variability also observed in NTL, the latter independently estimated from  $C_b$   
417 (Table 2). On average, NTL indicated that the oceanic plankton food webs analyzed had less than 3  
418 trophic levels, with maximum values in the Coastal and Trades biomes. These values were ca. one  
419 trophic level higher than the estimations of the trophic position for the largest plankton size class  
420 using  $\delta^{15}\text{N}$ . The difference would suggest that these planktonic food webs would support only one  
421 additional consumer level predating upon the largest size class analyzed. Some of consumers may  
422 well be the cnidaria and ctenophora removed from the largest size fraction of our samples, with  
423 estimated TP between 2.6 and 2.9 (Molina-Ramirez et al., 2015). The trophic position of large  
424 plankton and the PPMR values obtained in our study are indicative of feeding on small plankton by  
425 most consumers. This is expected in oligotrophic waters dominated by picophytoplankton, (Hunt  
426 et al., 2015), thus suggesting a common trophic structure related to the nutrient inputs. The resulting  
427 TTE values were in the range observed for pelagic (Hunt et al., 2015) but also benthic ecosystems  
428 (Jennings et al., 2001) indicating a relatively high efficiency in the transfer of energy through the  
429 food web. However, these values were lower than the average estimated of 70% for pelagic food  
430 webs (Zhou, 2006). For instance, using a global TTE average of 21% would increase the NTL to  
431 6.4, thus suggesting that the use of biomass spectra corrected by a TTE estimate from  $\delta^{15}\text{N}$  would  
432 take into account the larger number of trophic steps expected in oligotrophic waters (Sommer et al.,  
433 2002).

434 The low variability observed in  $\delta^{15}\text{N}$  slopes suggests that the overall trophic structure is similar  
435 across different planktonic communities, even when there are large variations in the nutrient sources  
436 and in the transmission of the energy through the food web. Organisms included in our spectra  
437 covered four orders of magnitude in individual size, implying a large difference in growth rates and  
438 biomass turnover times, and thus larger variability in  $\delta^{15}\text{N}$  composition of the small plankton versus  
439 large plankton (Jennings *et al.*, 2008). Such differences would explain the reported delays in the  
440 transmission of the  $\delta^{15}\text{N}$  signal from the N source (e.g. atmospheric N) to zooplankton predators  
441 (Rolff, 2000; Mompeán *et al.*, 2013). The normalization procedure applied in our study overcomes  
442 the time-lags and local variability of  $\delta^{15}\text{N}$  and, as normalized biomass spectra, provides a better tool  
443 to compare different communities than un-normalized spectra. Similarity in  $\delta^{15}\text{N}$  spectrum slopes  
444 and trophic structure across plankton communities can be expected if the underlying physiological  
445 processes are similar. Even when the nutrient (nitrogen) sources are different, the transmission up  
446 the food web is made by similar biochemical reactions for all planktonic communities. In the case  
447 of tropical and subtropical communities living in oligotrophic waters, recycling of dissolved organic  
448 matter through microbial communities is an important mechanism for transmission of nitrogen to

449 upper trophic levels (Mulholland, 2007). The homogeneity in trophic structure across pelagic  
450 ecosystems was already suggested by the first studies of biomass and abundance size spectra (Platt  
451 and Denman, 1978) and confirmed by subsequent studies including upper trophic levels (Hunt et  
452 al., 2015). Now, the application of biomass-normalized  $\delta^{15}\text{N}$  spectra to whole food webs, including  
453 fish and upper consumers, can be used to analyze their variability by considering nitrogen  
454 exchanges.

## 455 **CONCLUSIONS**

456 Biomass and  $\delta^{15}\text{N}$  size-spectra of plankton showed in general low variability, reflecting a large  
457 homogeneity in nitrogen sources and food web structure for the central oceans. The largest changes  
458 in spectral parameters and nitrogen sources were observed in the northwestern Atlantic and were  
459 related to inputs of atmospheric nitrogen, either from diazotrophic organisms or dust deposition.  
460 Mean predator-to-prey mass ratios between 4 and  $3 \cdot 10^3$  were estimated for all biomes, while mean  
461 trophic transfer efficiency values varied between 16% (coastal biome) and >20% (Trades and  
462 Westerlies biomes), thus suggesting that current estimates of the number of trophic levels based on  
463 a high average community efficiency may be too low, at least for most of the oligotrophic ocean.  
464 Both  $\delta^{15}\text{N}$  and biomass size-spectra suggest geographic homogeneity in the net transfer of nitrogen  
465 up the food web.

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## 478 **DATA ARCHIVING**

479 The biomass and stable isotope data used in this paper was deposited in the Malaspina-2010  
480 repository and can be accessed through the following links:

481 [http://metamalaspina.imedea.uib-csic.es:80/geonetwork?uuid=91b1175d-894c-4e48-980f-  
482 9cc325dcd79e](http://metamalaspina.imedea.uib-csic.es:80/geonetwork?uuid=91b1175d-894c-4e48-980f-9cc325dcd79e)

483 [http://metamalaspina.imedea.uib-csic.es:80/geonetwork?uuid=4f79814e-ebab-4e0f-aea9-  
484 25c88bacd25a](http://metamalaspina.imedea.uib-csic.es:80/geonetwork?uuid=4f79814e-ebab-4e0f-aea9-25c88bacd25a)

485

486 [Data from leg 8 are also available at the PANGAEA repository:](#)

487 <https://doi.pangaea.de/10.1594/PANGAEA.816451>

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628 **Table and Figure legends:**

629 Table 1. Mean±se values of the parameters of the size spectra for biomass ( $C_a$ : intercept;  $C_b$ : slope)  
630 and  $\delta^{15}\text{N}$  ( $\delta^{15}\text{N}_a$ : intercept;  $\delta^{15}\text{N}_b$ : slope) and of carbon biomass ( $C_{40-200}$ ,  $\text{mg C m}^{-3}$ ) and  $\delta^{15}\text{N}$   
631 ( $\delta^{15}\text{N}_{40-200}$ , ‰) for the 40-200  $\mu\text{m}$  size class in the three biomass sampled. Means not significantly  
632 different are marked in the same group (ANOVA and Dunnett-C test,  $P<0.05$ ). n: number of data.

633 Table 2. Mean±se values of the trophic position of the 2000-5000  $\mu\text{m}$  size-class ( $\text{TP}_{2000}$ ) and food  
634 web parameters derived from size spectra in the three biomass sampled. NTL: number of trophic  
635 levels (Zhou, 2006); PPMR: predator-to-prey mass ratio (Jennings et al., 2001); TTE: trophic transfer  
636 efficiency, as % (Barnes et al., 2010). Means not significantly different are marked in the same  
637 group (ANOVA and Dunnett-C test,  $P<0.05$ ). n: number of data.

638 Table 3. Median, minimum (min) and maximum (max) values of GWR coefficients for spectral  
639 parameters ( $C_a$ ,  $\delta^{15}\text{N}_a$  and  $\delta^{15}\text{N}_b$ ) estimated from logarithmic *Trichodesmium* abundance (Tricho,  
640  $\log_2$  trichomes  $\text{m}^{-3}$ ), depth of chlorophyll maximum (DCM, m) and mean monthly dust deposition ( $\text{g}$   
641  $\text{m}^{-2} \text{month}^{-1}$ ). All models were fit with  $P<0.05$  and include a constant term. Values from models for  
642  $C_b$  were not shown as GWR was not significant for this variable.

643 Table 1S. Parameters of the biomass-normalized biomass and  $\delta^{15}\text{N}$  size-spectra determined for  
644 plankton in the 0-200 m layer of stations of the Malaspina-2010 expedition. Lines in the form  $Y = a$   
645  $+ b \log_2(w)/\Delta w$  were fitted by least squares. w: mean individual size ( $\mu\text{g C}$ ) of organisms in each  
646 size-class;  $\Delta w$ : range of w for each size interval; r: correlation coefficient, se: standard error of the  
647 regression. All parameters were significant ( $P<0.05$ ).

648 Figure 1. Position of stations sampled during the Malaspina-2010 expedition for determination of  
649 size-fractionated plankton biomass and natural abundance of stable nitrogen isotopes. The colors  
650 indicate the biomes (black: Coastal, grey: Westerlies, white: Trades) according to Longhurst (2007).  
651 The arrows and numbers indicate the different cruise legs and track direction.

652 Figure 2. Examples of size spectra for plankton samples collected at two stations of leg 8 of the  
653 Malaspina-2010 expedition. a) un-normalized spectra for biomass ( $C$ ,  $\text{mg m}^{-3}$ ), b) un-normalized  
654 spectra for  $\delta^{15}\text{N}$  (‰), c)  $\delta^{15}\text{N}$  biomass-normalized spectra. Mean biomass of the size class ( $\log_2(w)$ )  
655 in  $\mu\text{g C}$ .

656 Figure 3. Distribution of plankton biomass (a,  $C_{40}$ ,  $\text{mg C m}^{-3}$ ) and  $\delta^{15}\text{N}$  (b,  $\delta^{15}\text{N}_{40}$  ‰) in the 40-200  
657  $\mu\text{m}$  size fraction for stations sampled during the Malaspina-2010 expedition. The vertical dashed

658 lines and numbers indicate the different cruise legs. The horizontal dashed line in b) indicates  
659  $\delta^{15}\text{N}=0$ . Note the inversion of values in the horizontal axis for leg 8.

660 Figure 4. Distribution of parameters of biomass (a, b) and  $\delta^{15}\text{N}$  (c, d) plankton size spectra  
661 computed for stations sampled during the Malaspina-2010 expedition. a), b): intercept of biomass  
662 ( $C_a$ ) and  $\delta^{15}\text{N}$  ( $\delta^{15}\text{N}_a$ ) spectra, respectively. c), d): slope of biomass ( $C_b$ ) and  $\delta^{15}\text{N}$  ( $\delta^{15}\text{N}_b$ ) spectra,  
663 respectively. Cruise legs indicated as in Fig. 3.

664 Figure 5. Distribution of *Trichodesmium* abundance (a, Tricho,  $\log(\text{trichomes}+1) \text{ m}^{-3}$ ), depth of the  
665 chlorophyll maximum (b, DCM, m) and mean dust deposition (c, MDU,  $\text{g m}^{-2}$ ) measured or  
666 estimated for stations sampled during the Malaspina-2010 expedition.

667 Figure 6. Vectors of variables projected in the space of the two main components of the PCA, .  
668 factors 1 and 2 accounting for 39% and 18% of total variance, respectively.  $C_a$  and  $C_b$ : intercept and  
669 slope of the biomass spectra, respectively;  $\delta^{15}\text{N}_a$  and  $\delta^{15}\text{N}_b$ : intercept and slope of the  $\delta^{15}\text{N}$  spectra,  
670 respectively; PP: primary production; Tricho: *Trichodesmium* abundance; MDU: mean dust  
671 deposition; MLD: mixed layer depth; DCM: depth of the chlorophyll maximum;  $C_{40}$  and  $\delta^{15}\text{N}_{40}$ :  
672 biomass and  $\delta^{15}\text{N}$  of the 40-200  $\mu\text{m}$  size-class, respectively; Chla: chlorophyll concentration;  $N_m^2$ :  
673 mean squared Brunt-Väisälä frequency.

674 Figure 7. Spatial autocorrelation (Moran's I) for spectral parameters estimated by GWR at different  
675 spatial scales. a), b):  $C_a$  and  $C_b$ : intercept and slope of the biomass spectra, respectively; c), d):  
676  $\delta^{15}\text{N}_a$  and  $\delta^{15}\text{N}_b$ : intercept and slope of the  $\delta^{15}\text{N}$  spectra, respectively.

677 Figure 1S. Relationship between the depth of the chlorophyll maximum (DCM, m) and the vertical  
678 flux of nitrate across the pycnocline by mechanical diffusion and salt fingers ( $\text{Flux}[\text{NO}_3]_{\text{tsf}}, \mu\text{mol}$   
679  $\text{m}^{-2} \text{d}^{-1}$ ). Nitrate fluxes from Fernández-Castro *et al.* (Fernández-Castro *et al.*, 2015).

680 Figure 2S. Variability of a) biomass ( $C, \text{mg m}^{-3}$ ) and b)  $\delta^{15}\text{N}$  values (‰) with mean organism size  
681  $[\log_2(w)]$ . Samples for all stations were plotted together.

682 Figure 3S. Distribution of (a, b) intercept ( $C_a, \delta^{15}\text{N}_a$ ) and (c, d) slope ( $C_b, \delta^{15}\text{N}_b$ ) of biomass and  
683  $\delta^{15}\text{N}$  plankton size spectra, respectively, e) *Trichodesmium* abundance (Tricho,  $\log \text{trichomes m}^{-3}$ ),  
684 (f) primary production ( $\text{mg C m}^{-2} \text{d}^{-1}$ ), g) mean dust deposition (MDU,  $\text{g m}^{-2}$ ), h) depth of the  
685 chlorophyll maximum (DCM, m), i) integrated chlorophyll-a concentration (Chla,  $\text{mg m}^{-2}$ ), j) mixed  
686 layer depth (MLD, m), k) mean squared Brunt-Väisälä frequency ( $N_m^2, 10^{-4} \text{ s}^{-2}$ ), and l) standard  
687 deviation of the squared Brunt-Väisälä frequency ( $N_m^2 \text{sd}, 10^{-4} \text{ s}^{-2}$ ) measured or estimated for stations  
688 sampled during the Malaspina-2010 expedition.

689 Figure 4S. Relationship between the slope of the isotope spectrum ( $\delta^{15}\text{N}_b$ ) and the trophic position  
690 of plankton in the size fraction 2000-5000  $\mu\text{m}$  ( $\text{TP}_{2000}$ ). Trophic position computed as  
691  $\text{TP}_{2000} = \text{TP}_{\text{base}} + (\delta^{15}\text{N}_{2000} - \delta^{15}\text{N}_{\text{base}}) / \Delta^{15}\text{N}$ , where  $\text{TP}_{\text{base}}$  is the trophic position of the baseline (1.5,  
692 estimated as a mixture of phyto- and zooplankton in the size-fraction 40-200  $\mu\text{m}$ ),  $\delta^{15}\text{N}_{\text{base}}$  is the  
693 measured  $\delta^{15}\text{N}$  in the 40-200  $\mu\text{m}$  size-fraction and  $\Delta^{15}\text{N}$  the average increase in  $\delta^{15}\text{N}$  between  
694 trophic levels (e.g. Post, 2002). The white dots represent outliers ( $\delta^{15}\text{N}_b > -0.7$ ) and were not used in  
695 the regression.



696 **Tables**

697 Table 1. Mean±se values of the parameters of the size spectra for biomass ( $C_a$ : intercept;  $C_b$ : slope)  
 698 and  $\delta^{15}\text{N}$  ( $\delta^{15}\text{N}_a$ : intercept;  $\delta^{15}\text{N}_b$ : slope) and of carbon biomass ( $C_{40-200}$ , mg C m<sup>-3</sup>) and  $\delta^{15}\text{N}$   
 699 ( $\delta^{15}\text{N}_{40-200}$ , ‰) for the 40-200  $\mu\text{m}$  size class in the three biomass sampled. Means not significantly  
 700 different are marked in the same group (ANOVA and Dunnett-C test,  $P < 0.05$ ). n: number of data.

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variable	biome	mean	se	n	group
$C_a$	Coastal	8.840	0.166	11	a
	Westerlies	7.954	0.095	55	b
	Trades	8.961	0.071	79	a
$C_b$	Coastal	-0.871	0.021	11	a
	Westerlies	-0.941	0.009	55	b
	Trades	-0.884	0.006	79	a
$\delta^{15}\text{N}_a$	Coastal	10.589	0.230	11	a
	Westerlies	9.444	0.163	55	b
	Trades	10.444	0.183	79	a
$\delta^{15}\text{N}_b$	Coastal	-0.867	0.008	11	a
	Westerlies	-0.824	0.010	55	a
	Trades	-0.832	0.012	79	a
$C_{40}$	Coastal	1.314	0.193	11	a
	Westerlies	1.083	0.068	55	a
	Trades	1.491	0.090	79	a
$\delta^{15}\text{N}_{40}$	Coastal	5.520	0.881	11	a
	Westerlies	2.723	0.422	55	b
	Trades	5.810	0.365	79	a

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709 Table 2. Mean±se values of the trophic position of the 2000-5000  $\mu\text{m}$  size-class (TP<sub>2000</sub>) and food  
710 web parameters derived from size spectra in the three biomass sampled. NTL: number of trophic  
711 levels (Zhou, 2006); PPMR: predator-to-prey mass ratio (Jennings et al., 2001); TTE: trophic  
712 transfer efficiency, as % (Barnes et al., 2010). Means not significantly different are marked in the  
713 same group (ANOVA and Dunnett-C test, P<0.05). n: number of data.

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variable	biome	mean	se	n	group
TP <sub>2000</sub>	Coastal	1.47	0.12	11	b
	Westerlies	1.91	0.07	55	a
	Trades	1.95	0.07	79	a
NTL	Coastal	2.81	0.07	11	a
	Westerlies	2.60	0.03	55	b
	Trades	2.76	0.02	79	a
PPMR	Coastal	696.71	199.75	11	a
	Westerlies	378.49	148.22	55	a
	Trades	723.12	203.19	79	a
TTE	Coastal	16.44	1.55	11	b
	Westerlies	24.91	1.72	55	a
	Trades	21.66	1.42	79	a, b

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Table 3. Median, minimum (min) and maximum (max) values of GWR coefficients for spectral parameters ( $C_a$ ,  $\delta^{15}N_a$  and  $\delta^{15}N_b$ ) estimated from logarithmic *Trichodesmium* abundance (Tricho, log trichomes  $m^{-3}$ ), depth of chlorophyll maximum (DCM, m) and mean monthly dust deposition ( $g\ m^{-2}\ month^{-1}$ ). All models were fit with  $P < 0.05$  and include a constant term. Values from models for  $C_b$  are not shown as GWR was not significant for this variable.

Dependent	constant			Tricho			DCM			MDU		
	median	min	max	median	min	max	median	min	max	median	min	max
$C_a$	9.162	5.842	12.391	0.001	-0.036	0.058	-0.002	-0.032	0.004	0.327	-5.685	7.134
$\delta^{15}N_a$	11.194	7.444	15.799	-0.004	-0.076	0.063	-0.002	-0.045	0.029	-2.416	-20.418	3.535
$\delta^{15}N_b$	-0.879	-1.473	-0.722	0.000	-0.003	0.004	0.000	-0.001	0.004	0.089	-0.088	0.682

## Supplement

Table 1S. Parameters of the biomass-normalized biomass and  $\delta^{15}\text{N}$  size-spectra determined for plankton in the 0-200 m layer of stations of the Malaspina-2010 expedition. Lines in the form  $Y = a + b \log_2(w)/\Delta w$  were fitted by least squares.  $w$ : mean individual size ( $\mu\text{g C}$ ) of organisms in each size-class;  $\Delta w$ : range of  $w$  for each size interval;  $r$ : correlation coefficient,  $se$ : standard error of the regression. All parameters were significant ( $P < 0.05$ ).

leg	station	date	longitude	latitude	biome	biomass spectra				$\delta^{15}\text{N}$ spectra			
						a	b	r	se	a	b	r	se
1	10	26/12/2010	-26.0016	14.5195	Trades	8.648	-0.847	-0.995	0.426	10.401	-0.804	-0.997	0.293
1	13	29/12/2010	-25.9957	7.3283	Trades	9.021	-0.874	-0.995	0.459	10.047	-0.830	-0.990	0.606
1	15	31/12/2010	-26.0169	2.4625	Trades	7.858	-0.752	-0.981	0.753	10.840	-0.830	-0.993	0.489
1	17	02/01/2011	-27.3265	-3.0306	Trades	8.877	-0.761	-0.979	0.813	10.874	-0.830	-0.993	0.506
2	35	27/01/2011	-11.8002	-28.6088	Trades	7.952	-0.844	-0.953	1.373	10.971	-0.893	-0.992	0.566
2	39	31/01/2011	0.9678	-30.8815	Trades	9.379	-0.938	-0.972	1.146	11.352	-0.829	-0.995	0.430
3	45	13/02/2011	25.5633	-35.1366	Coastal	9.100	-0.756	-0.996	0.345	11.171	-0.899	-0.991	0.622
3	46	14/02/2011	27.5456	-34.8371	Coastal	9.259	-0.953	-0.986	0.808	10.666	-0.874	-0.990	0.634
3	47	15/02/2011	31.0858	-34.4448	Coastal	8.786	-0.899	-0.991	0.606	10.597	-0.852	-0.994	0.471
3	48	16/02/2011	33.7251	-34.1728	Trades	8.529	-0.924	-0.966	1.263	10.550	-0.896	-0.987	0.752
3	49	17/02/2011	37.0018	-33.8828	Trades	8.606	-0.808	-0.992	0.532	10.702	-0.864	-0.993	0.535
3	50	18/02/2011	39.8805	-33.5322	Trades	7.522	-0.999	-0.996	0.467	10.625	-0.913	-0.993	0.548
3	51	19/02/2011	43.2477	-33.1928	Trades	8.651	-0.882	-0.986	0.773	10.626	-0.899	-0.988	0.707
3	52	23/02/2011	63.2474	-27.9776	Trades	8.826	-0.846	-0.991	0.593	10.521	-0.855	-0.999	0.233
3	53	25/02/2011	66.4931	-28.1283	Trades	9.123	-0.837	-0.990	0.595	10.225	-0.864	-0.994	0.483
3	54	26/02/2011	69.4122	-29.3618	Trades	8.768	-0.821	-0.991	0.560	10.653	-0.860	-0.995	0.443
3	55	27/02/2011	72.4499	-29.5619	Trades	8.495	-0.899	-0.971	1.130	10.651	-0.875	-0.995	0.429
3	56	28/02/2011	76.0806	-29.9046	Trades	8.243	-0.912	-0.967	1.224	11.080	-0.866	-0.993	0.530
3	57	01/03/2011	79.6065	-29.8342	Trades	8.741	-0.861	-0.986	0.743	10.889	-0.893	-0.990	0.639
3	58	02/03/2011	82.6240	-29.8106	Trades	8.614	-0.853	-0.989	0.658	11.313	-0.856	-0.995	0.456
3	59	03/03/2011	86.2619	-29.7575	Trades	8.298	-0.973	-0.979	1.027	11.508	-0.856	-0.994	0.489
3	60	04/03/2011	89.4798	-29.6831	Trades	8.762	-0.871	-0.982	0.864	11.169	-0.866	-0.990	0.638
3	61	05/03/2011	92.9838	-29.6423	Trades	8.681	-0.783	-0.994	0.443	11.076	-0.870	-0.991	0.597
3	62	06/03/2011	96.3948	-29.5818	Trades	8.931	-0.798	-0.985	0.718	11.128	-0.864	-0.992	0.564

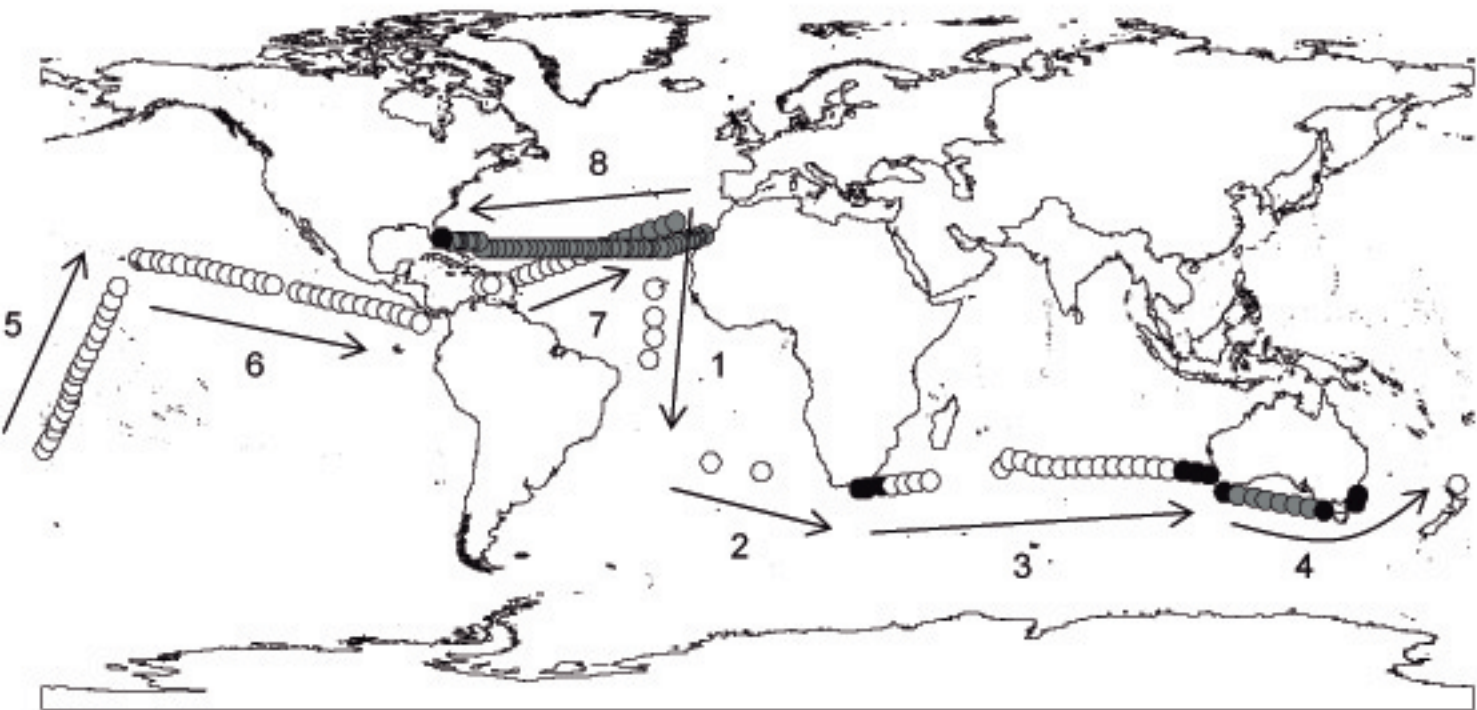
3	63	07/03/2011	99.9978	-29.9066	Trades	8.661	-0.878	-0.984	0.816	11.443	-0.892	-0.993	0.553
3	64	08/03/2011	103.3091	-30.3329	Trades	8.463	-0.885	-0.982	0.880	10.820	-0.851	-0.991	0.578
3	65	09/03/2011	107.2080	-30.8139	Trades	8.800	-0.857	-0.984	0.785	10.294	-0.916	-0.993	0.576
3	66	10/03/2011	110.1896	-31.1453	Coastal	9.240	-0.840	-0.982	0.828	9.554	-0.871	-0.994	0.481
3	67	11/03/2011	113.4476	-31.5598	Coastal	8.529	-0.884	-0.965	1.227	10.195	-0.857	-0.984	0.786
3	68	12/03/2011	113.3935	-31.5493	Coastal	9.933	-0.836	-0.995	0.420	9.999	-0.878	-0.991	0.620
4	69	18/03/2011	120.8448	-36.6445	Coastal	8.216	-0.900	-0.950	1.515	11.232	-0.883	-0.992	0.574
4	70	19/03/2011	124.8610	-37.3508	Westerlies	8.536	-0.909	-0.981	0.927	11.327	-0.911	-0.988	0.728
4	71	20/03/2011	127.7512	-37.8550	Westerlies	9.244	-0.886	-0.986	0.757	11.692	-0.873	-0.991	0.619
4	72	21/03/2011	131.5408	-38.5484	Westerlies	8.957	-0.931	-0.983	0.880	12.015	-0.874	-0.992	0.564
4	73	22/03/2011	135.2284	-39.2225	Westerlies	9.408	-0.932	-0.990	0.668	11.912	-0.872	-0.992	0.558
4	74	23/03/2011	138.7843	-39.8431	Westerlies	9.016	-0.820	-0.981	0.835	11.969	-0.854	-0.993	0.527
4	75	24/03/2011	142.5004	-40.5312	Westerlies	8.595	-0.965	-0.988	0.765	11.915	-0.886	-0.992	0.560
4	76	25/03/2011	146.5807	-39.2681	Coastal	9.077	-0.776	-0.987	0.638	11.959	-0.888	-0.993	0.533
4	77	27/03/2011	150.4144	-38.6609	Coastal	8.207	-0.829	-0.979	0.888	10.974	-0.884	-0.991	0.613
4	78	28/03/2011	150.9914	-36.6807	Coastal	8.749	-0.979	-0.987	0.816	10.990	-0.856	-0.992	0.562
5	79	17/04/2011	176.0029	-34.0518	Trades	8.848	-0.864	-0.985	0.768	11.143	-0.868	-0.995	0.462
5	82	19/04/2011	-179.5199	-25.4835	Trades	8.915	-0.953	-0.986	0.811	9.772	-0.882	-0.990	0.648
5	83	20/04/2011	-178.2187	-23.3642	Trades	9.081	-0.892	-0.985	0.805	9.160	-0.921	-0.986	0.798
5	84	21/04/2011	-176.9193	-20.6623	Trades	9.067	-0.907	-0.996	0.392	8.698	-0.799	-0.982	0.788
5	85	22/04/2011	-175.8323	-18.5553	Trades	8.476	-0.925	-0.996	0.432	9.583	-0.895	-0.983	0.857
5	86	23/04/2011	-174.4879	-15.9029	Trades	8.556	-0.917	-0.982	0.909	11.151	-0.886	-0.991	0.599
5	87	24/04/2011	-173.3727	-13.5315	Trades	8.525	-1.023	-0.977	1.140	11.617	-0.891	-0.989	0.695
5	88	25/04/2011	-172.6531	-11.2152	Trades	8.601	-0.860	-0.978	0.948	11.995	-0.861	-0.992	0.550
5	89	26/04/2011	-172.3362	-9.4545	Trades	8.593	-0.900	-0.986	0.774	12.101	-0.861	-0.992	0.568
5	90	27/04/2011	-171.4322	-7.0484	Trades	9.987	-0.928	-0.988	0.734	11.115	-0.858	-0.992	0.566
5	91	28/04/2011	-170.7707	-5.7366	Trades	8.445	-0.812	-0.965	1.127	10.261	-0.854	-0.987	0.702
5	92	29/04/2011	-169.4625	-3.4106	Trades	9.120	-0.905	-0.988	0.730	10.176	-0.864	-0.991	0.587
5	93	30/04/2011	-168.3567	-1.3037	Trades	9.710	-0.855	-0.984	0.783	10.590	-0.842	-0.993	0.503
5	94	01/05/2011	-166.8465	1.6188	Trades	10.179	-0.889	-0.986	0.758	10.889	-0.853	-0.992	0.539
5	95	02/05/2011	-165.7639	3.7850	Trades	10.878	-0.892	-0.991	0.600	10.964	-0.850	-0.992	0.540
5	96	03/05/2011	-164.4105	6.9823	Trades	9.659	-0.825	-0.984	0.766	11.857	-0.869	-0.992	0.582
5	97	04/05/2011	-163.5270	9.2177	Trades	8.987	-0.828	-0.984	0.776	11.913	-0.849	-0.995	0.440
5	98	05/05/2011	-162.4108	11.5919	Trades	7.868	-0.835	-0.975	0.977	11.242	-0.825	-0.994	0.455
5	99	06/05/2011	-160.8179	14.9963	Trades	8.716	-0.841	-0.984	0.768	11.215	-0.867	-0.990	0.629
6	101	14/05/2011	-155.6640	21.8908	Trades	7.891	-0.923	-0.954	1.487	10.190	-0.869	-0.988	0.684
6	102	15/05/2011	-153.4199	21.5712	Trades	8.774	-0.857	-0.989	0.669	9.903	-0.870	-0.983	0.836

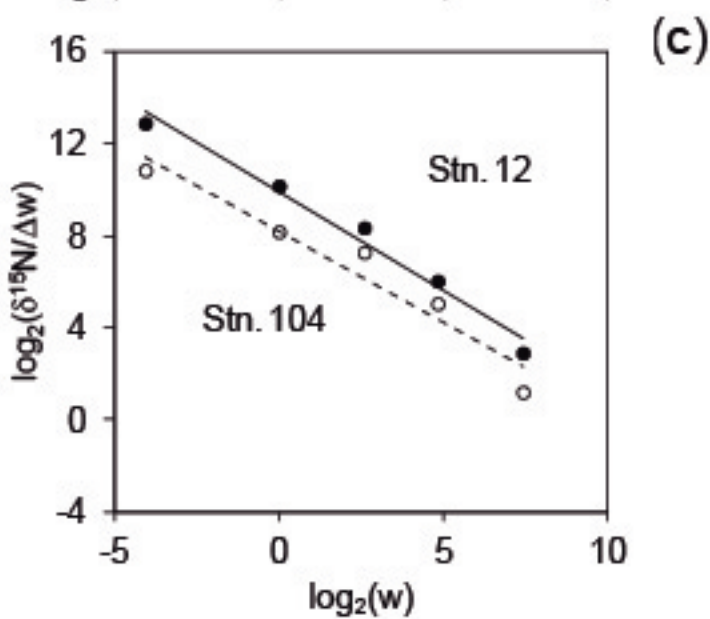
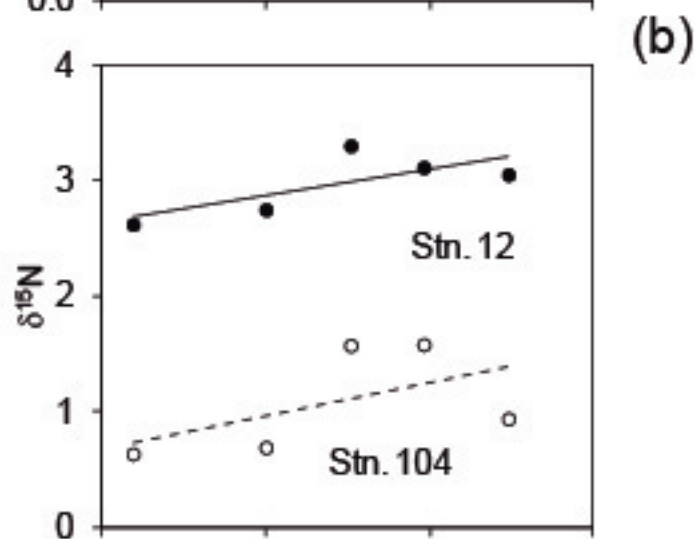
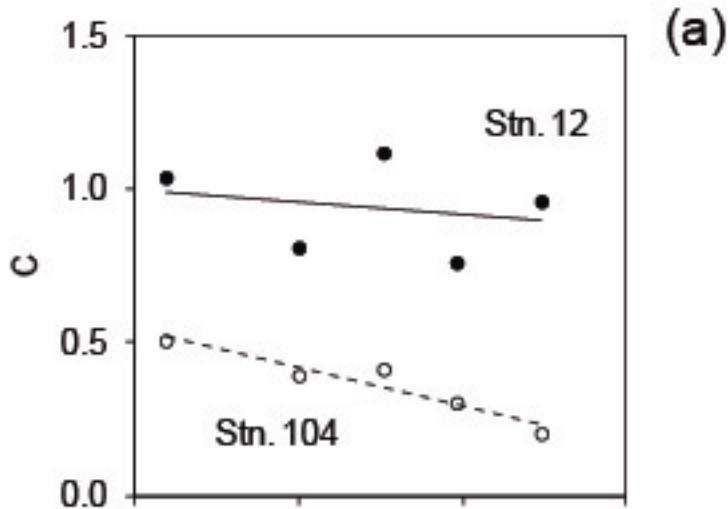
6	103	16/05/2011	-150.3591	21.0660	Trades	8.947	-0.848	-0.989	0.644	11.371	-0.868	-0.994	0.498
6	104	17/05/2011	-148.3402	20.8025	Trades	7.580	-0.872	-0.986	0.764	11.193	-0.862	-0.993	0.525
6	105	19/05/2011	-145.2160	20.3406	Trades	8.886	-0.808	-0.992	0.538	10.278	-0.806	-0.995	0.398
6	106	20/05/2011	-141.6154	19.8989	Trades	8.320	-0.824	-0.993	0.503	11.293	-0.859	-0.994	0.492
6	107	21/05/2011	-138.9657	19.2792	Trades	8.751	-0.840	-0.990	0.626	11.496	-0.873	-0.992	0.550
6	108	22/05/2011	-136.1774	18.6483	Trades	9.273	-0.908	-0.995	0.448	11.678	-0.849	-0.991	0.600
6	109	23/05/2011	-133.2623	18.0538	Trades	9.282	-0.866	-0.984	0.791	11.381	-0.867	-0.992	0.578
6	110	24/05/2011	-130.5909	17.4013	Trades	9.592	-0.951	-0.988	0.752	11.418	-0.862	-0.993	0.507
6	111	25/05/2011	-127.5755	16.6239	Trades	9.384	-0.887	-0.977	0.988	11.692	-0.860	-0.991	0.604
6	112	26/05/2011	-124.5040	15.9132	Trades	8.919	-0.781	-0.979	0.832	11.780	-0.847	-0.994	0.464
6	113	27/05/2011	-121.9954	15.3108	Trades	9.244	-0.880	-0.976	1.003	11.750	-0.868	-0.992	0.567
6	115	28/05/2011	-115.7725	13.7653	Trades	9.673	-0.932	-0.971	1.168	11.736	-0.857	-0.994	0.465
6	116	29/05/2011	-113.2701	13.1952	Trades	9.160	-0.952	-0.973	1.158	11.398	-0.838	-0.988	0.659
6	117	30/05/2011	-110.3917	12.4943	Trades	9.301	-0.908	-0.985	0.809	11.633	-0.865	-0.993	0.513
6	118	31/05/2011	-108.0495	11.9936	Trades	9.401	-0.949	-0.974	1.136	11.736	-0.857	-0.992	0.543
6	119	01/06/2011	-105.0116	11.3552	Trades	9.233	-0.861	-0.980	0.901	11.599	-0.857	-0.994	0.465
6	120	02/06/2011	-102.4479	10.7564	Trades	9.956	-0.872	-0.981	0.870	11.616	-0.874	-0.991	0.590
6	121	03/06/2011	-99.2499	10.0847	Trades	9.344	-0.906	-0.984	0.847	11.748	-0.854	-0.993	0.510
6	122	04/06/2011	-96.3360	9.4393	Trades	9.951	-0.821	-0.992	0.531	11.726	-0.863	-0.994	0.484
6	123	05/06/2011	-93.1477	8.7615	Trades	9.963	-0.810	-0.955	1.289	11.067	-0.864	-0.993	0.528
6	124	06/06/2011	-90.3620	8.1413	Trades	9.415	-0.936	-0.968	1.230	10.769	-0.838	-0.992	0.556
6	125	07/06/2011	-87.9549	7.2239	Trades	10.591	-0.887	-0.978	0.976	10.694	-0.852	-0.990	0.611
6	126	08/06/2011	-84.7942	5.9111	Trades	9.030	-0.888	-0.984	0.821	11.304	-0.849	-0.994	0.481
7	129	22/06/2011	-69.2892	15.0732	Trades	9.560	-0.905	-0.969	1.183	7.679	-0.629	-0.983	0.601
7	130	23/06/2011	-67.0695	15.5763	Trades	10.316	-0.962	-0.958	1.474	9.002	-0.910	-0.975	1.067
7	131	25/06/2011	-59.8288	17.4276	Trades	8.522	-0.895	-0.990	0.637	9.125	-0.823	-0.988	0.663
7	132	26/06/2011	-57.8094	18.0651	Trades	9.067	-0.913	-0.985	0.807	5.091	-0.404	-0.600	2.749
7	133	27/06/2011	-55.1573	19.0153	Trades	8.855	-0.914	-0.983	0.868	5.977	-0.597	-0.959	0.906
7	134	28/06/2011	-52.6289	20.0110	Trades	9.137	-0.928	-0.987	0.774	3.766	-0.237	-0.847	0.758
7	135	29/06/2011	-50.1433	20.8033	Trades	8.459	-1.001	-0.991	0.708	7.210	-0.738	-0.988	0.578
7	136	30/06/2011	-47.7887	21.7383	Trades	8.809	-0.976	-0.979	1.028	6.821	-0.776	-0.783	3.148
7	137	01/07/2011	-44.5286	22.8634	Trades	8.803	-0.966	-0.982	0.948	6.207	-0.463	-0.990	0.343
7	138	02/07/2011	-41.9070	23.7336	Trades	8.835	-0.975	-0.991	0.684	8.202	-0.797	-0.966	1.083
7	139	03/07/2011	-38.7107	24.8569	Trades	9.015	-0.892	-0.989	0.667	8.371	-0.773	-0.988	0.618
7	140	04/07/2011	-35.2705	26.1084	Westerlies	8.413	-0.990	-0.977	1.100	8.924	-0.815	-0.992	0.520
7	141	05/07/2011	-32.8858	26.9174	Westerlies	8.511	-0.897	-0.994	0.515	9.459	-0.808	-0.989	0.624
7	142	06/07/2011	-29.6718	27.9803	Westerlies	8.224	-0.947	-0.987	0.783	9.737	-0.873	-0.993	0.536

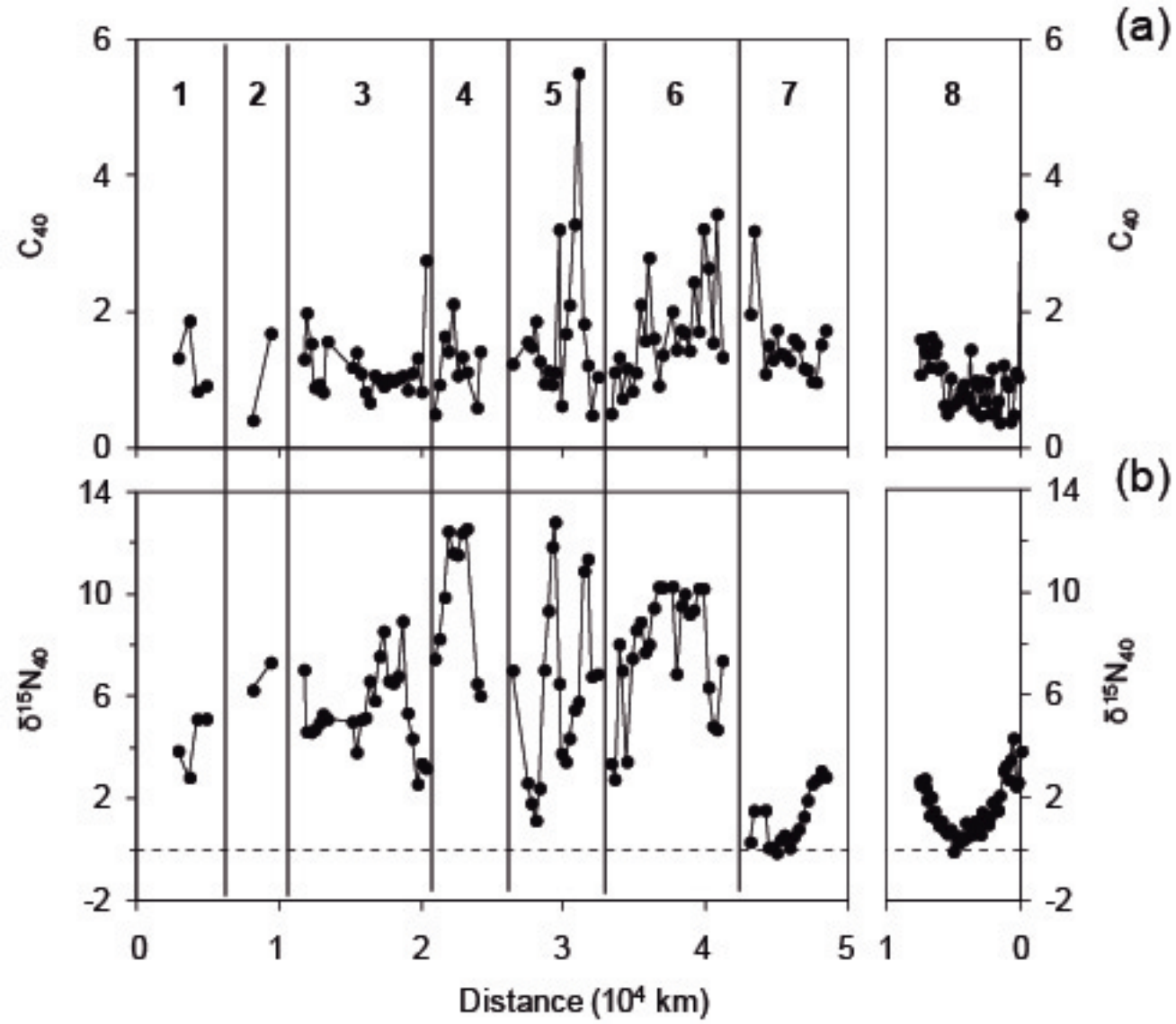
7	143	07/07/2011	-26.9515	28.8747	Westerlies	8.375	-0.877	-0.992	0.583	9.758	-0.868	-0.995	0.457
7	144	08/07/2011	-23.6907	29.9667	Westerlies	9.032	-0.940	-0.985	0.836	10.052	-0.839	-0.994	0.470
7	145	09/07/2011	-20.6366	30.9602	Westerlies	9.128	-0.960	-0.985	0.846	9.980	-0.823	-0.992	0.518
8	001	28/01/2011	-13.3418	27.7750	Westerlies	10.346	-0.749	-0.990	0.552	10.549	-0.829	-0.987	0.683
8	012	29/01/2011	-15.3478	27.1285	Westerlies	8.298	-0.891	-0.992	0.573	9.893	-0.855	-0.989	0.643
8	016	30/01/2011	-16.7797	26.6630	Westerlies	7.898	-0.905	-0.998	0.295	9.845	-0.862	-0.986	0.754
8	020	31/01/2011	-18.2703	26.1837	Westerlies	7.675	-0.849	-0.979	0.903	10.572	-0.877	-0.990	0.650
8	024	01/02/2011	-20.0092	25.6257	Westerlies	7.336	-0.785	-0.994	0.430	10.342	-0.868	-0.987	0.728
8	027	02/02/2011	-21.4565	25.1563	Westerlies	7.730	-0.955	-0.987	0.794	9.954	-0.846	-0.993	0.518
8	030	03/02/2011	-22.9763	24.6612	Westerlies	7.821	-0.916	-0.998	0.257	10.275	-0.827	-0.992	0.534
8	033	04/02/2011	-24.7080	24.5012	Westerlies	7.943	-0.971	-0.994	0.540	10.161	-0.860	-0.983	0.827
8	037	05/02/2011	-27.1497	24.5012	Westerlies	7.214	-0.807	-0.990	0.573	9.523	-0.871	-0.992	0.563
8	039	06/02/2011	-28.3718	24.5010	Westerlies	7.614	-0.892	-0.993	0.552	9.166	-0.837	-0.993	0.520
8	040	07/02/2011	-28.9843	24.4990	Westerlies	7.179	-0.963	-0.991	0.648	9.302	-0.827	-0.991	0.561
8	043	08/02/2011	-30.8183	24.4992	Westerlies	7.167	-0.928	-0.982	0.906	9.537	-0.826	-0.992	0.547
8	045	09/02/2011	-32.0403	24.5010	Westerlies	7.621	-0.936	-0.996	0.420	9.113	-0.946	-0.991	0.667
8	048	10/02/2011	-33.8727	24.5008	Westerlies	6.927	-0.890	-0.986	0.755	9.311	-0.739	-0.992	0.482
8	052	11/02/2011	-34.7195	24.4998	Westerlies	7.459	-1.004	-0.995	0.507	8.613	-0.835	-0.984	0.774
8	055	12/02/2011	-37.1078	24.5017	Westerlies	7.542	-0.901	-0.994	0.494	8.392	-0.957	-0.982	0.941
8	059	13/02/2011	-38.9537	24.4998	Westerlies	7.670	-0.963	-0.999	0.264	8.942	-0.843	-0.996	0.366
8	061	14/02/2011	-39.8760	24.5022	Westerlies	7.480	-0.768	-0.995	0.388	8.218	-0.686	-0.995	0.336
8	064	15/02/2011	-41.2633	24.4998	Westerlies	7.567	-0.994	-0.995	0.494	8.746	-0.784	-0.983	0.748
8	068	16/02/2011	-43.1097	24.5002	Westerlies	7.578	-1.000	-0.998	0.339	8.838	-0.788	-0.994	0.458
8	071	17/02/2011	-44.4953	24.4998	Westerlies	7.388	-0.883	-0.994	0.489	8.390	-0.872	-0.953	1.410
8	075	18/02/2011	-46.3437	24.5000	Westerlies	7.932	-1.023	-0.996	0.457	8.200	-0.745	-0.981	0.761
8	078	19/02/2011	-47.7278	24.5007	Westerlies	7.666	-0.891	-0.993	0.551	7.924	-0.820	-0.979	0.866
8	082	20/02/2011	-49.5755	24.5000	Westerlies	7.103	-1.075	-0.992	0.709	8.676	-0.791	-0.988	0.632
8	085	21/02/2011	-50.9612	24.5002	Westerlies	7.647	-0.929	-0.985	0.821	8.209	-0.721	-0.964	1.013
8	088	22/02/2011	-52.3467	24.5000	Westerlies	7.465	-0.991	-0.998	0.310	8.061	-0.938	-0.944	1.678
8	091	23/02/2011	-54.0342	24.4865	Westerlies	7.224	-0.963	-0.997	0.377	7.622	-0.602	-0.978	0.660
8	094	24/02/2011	-55.8743	24.5005	Westerlies	7.109	-0.970	-0.993	0.587	7.297	-0.982	-0.829	3.384
8	097	25/02/2011	-57.7157	24.5008	Westerlies	7.164	-0.971	-0.994	0.528	6.013	-0.561	-0.582	4.002
8	100	26/02/2011	-59.5530	24.5000	Westerlies	7.556	-1.045	-0.995	0.510	8.878	-0.728	-0.978	0.787
8	104	27/02/2011	-62.0047	24.5012	Westerlies	7.079	-0.984	-0.988	0.772	8.200	-0.792	-0.968	1.052
8	107	28/02/2011	-63.8430	24.5015	Westerlies	7.165	-0.991	-0.993	0.611	8.391	-0.703	-0.997	0.282
8	110	01/03/2011	-65.6820	24.5007	Westerlies	8.001	-0.984	-0.993	0.612	8.770	-0.804	-0.987	0.658
8	113	02/03/2011	-67.5202	24.4997	Westerlies	7.807	-1.025	-0.996	0.461	8.773	-0.734	-0.994	0.423

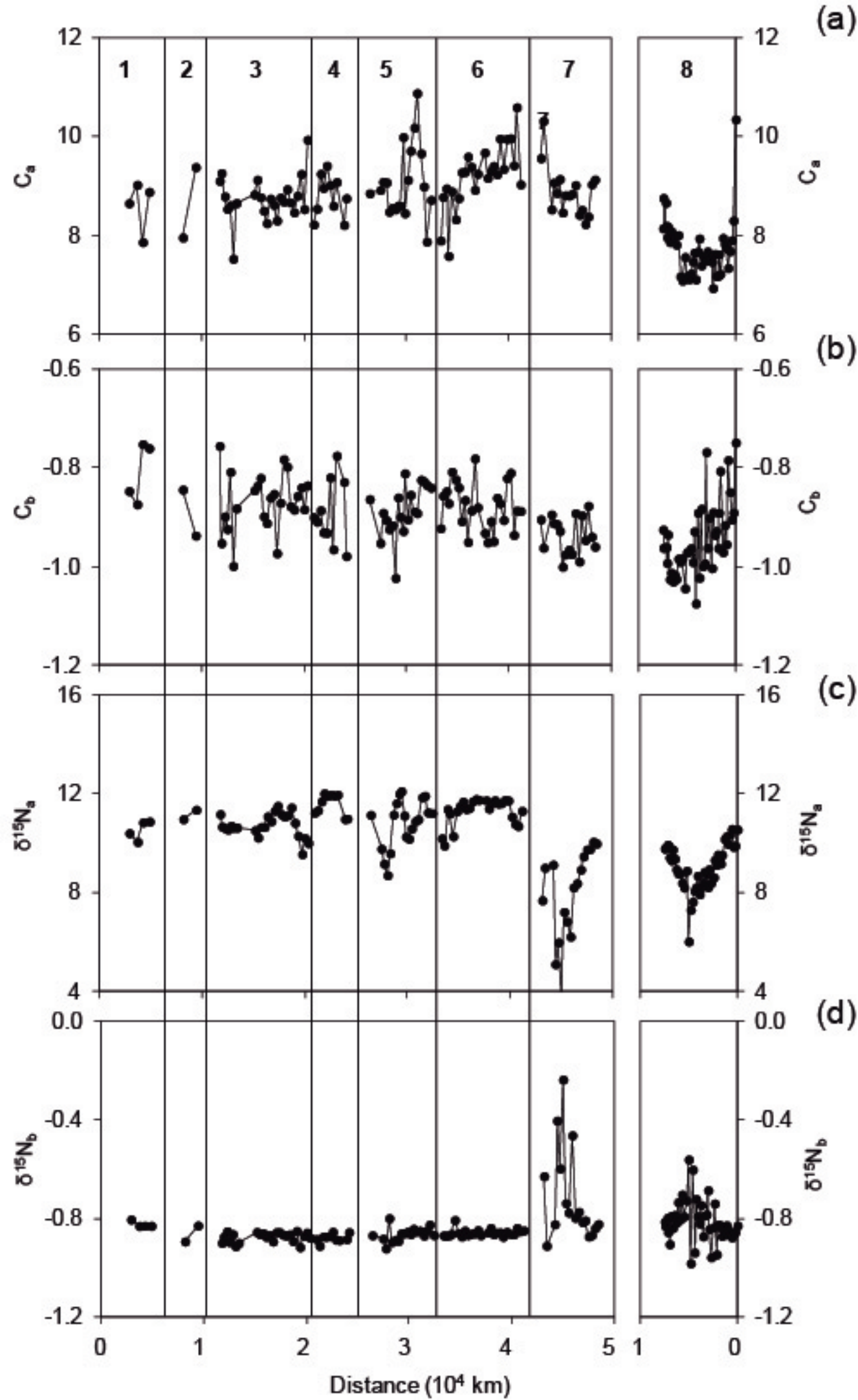
8	116	03/03/2011	-69.1327	24.5007	Westerlies	7.992	-1.017	-0.998	0.287	8.934	-0.817	-0.986	0.703
8	120	05/03/2011	-70.3220	26.2038	Westerlies	7.990	-1.031	-0.995	0.525	9.364	-0.786	-0.991	0.545
8	125	06/03/2011	-72.2647	26.2032	Westerlies	8.083	-1.013	-1.000	0.047	9.738	-0.839	-0.987	0.707
8	128	07/03/2011	-73.4317	26.2035	Westerlies	7.862	-1.026	-0.997	0.387	9.297	-0.789	-0.986	0.689
8	133	08/03/2011	-75.1557	26.2038	Westerlies	8.165	-0.936	-0.998	0.280	9.412	-0.904	-0.980	0.936
8	137	09/03/2011	-75.8622	26.2042	Westerlies	7.957	-0.993	-0.998	0.297	9.855	-0.799	-0.995	0.411
8	143	10/03/2011	-76.6218	26.2037	Westerlies	8.192	-0.961	-0.999	0.258	9.899	-0.856	-0.992	0.560
8	149	11/03/2011	-76.9525	26.2030	Westerlies	8.656	-0.961	-0.995	0.496	9.928	-0.797	-0.995	0.426
8	157	13/03/2011	-79.3063	27.0003	Westerlies	8.753	-0.962	-0.993	0.590	9.781	-0.830	-0.994	0.447
8	161	13/03/2011	-79.5658	27.0012	Coastal	8.138	-0.926	-0.997	0.349	9.784	-0.813	-0.996	0.385

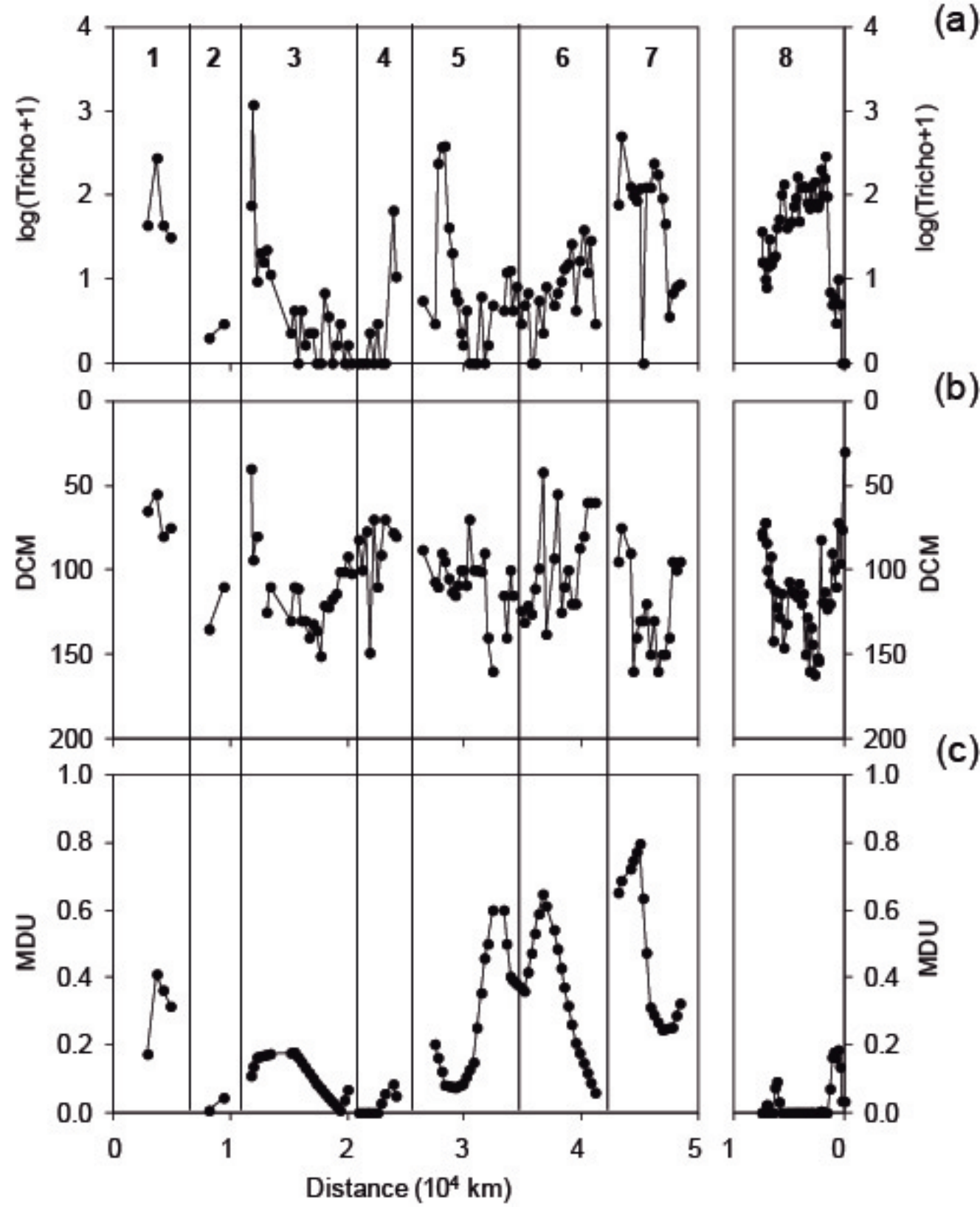


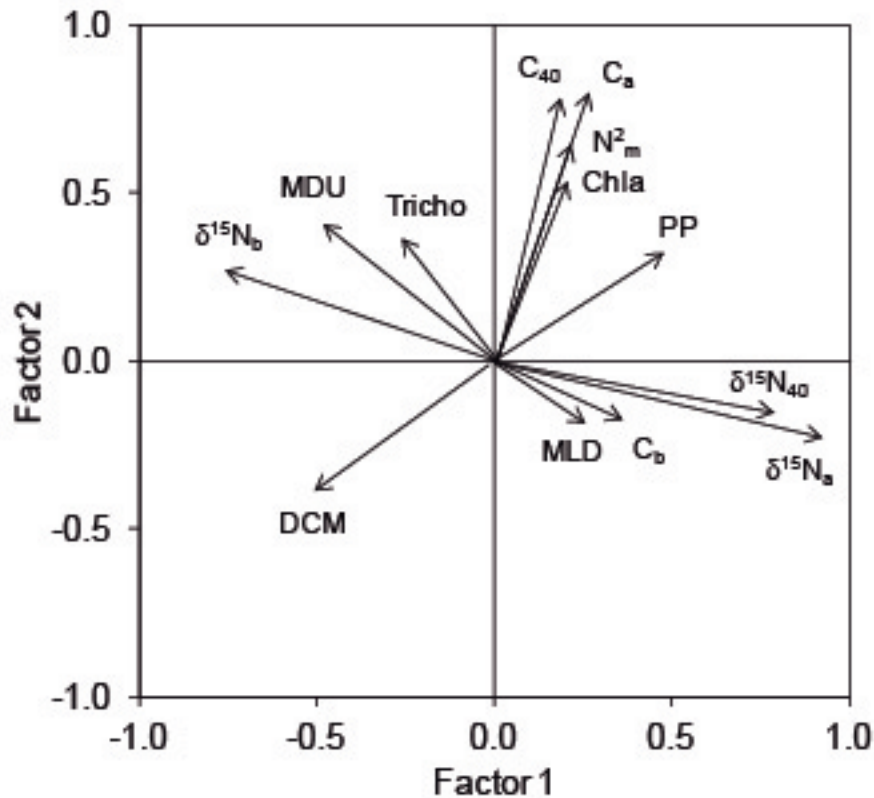


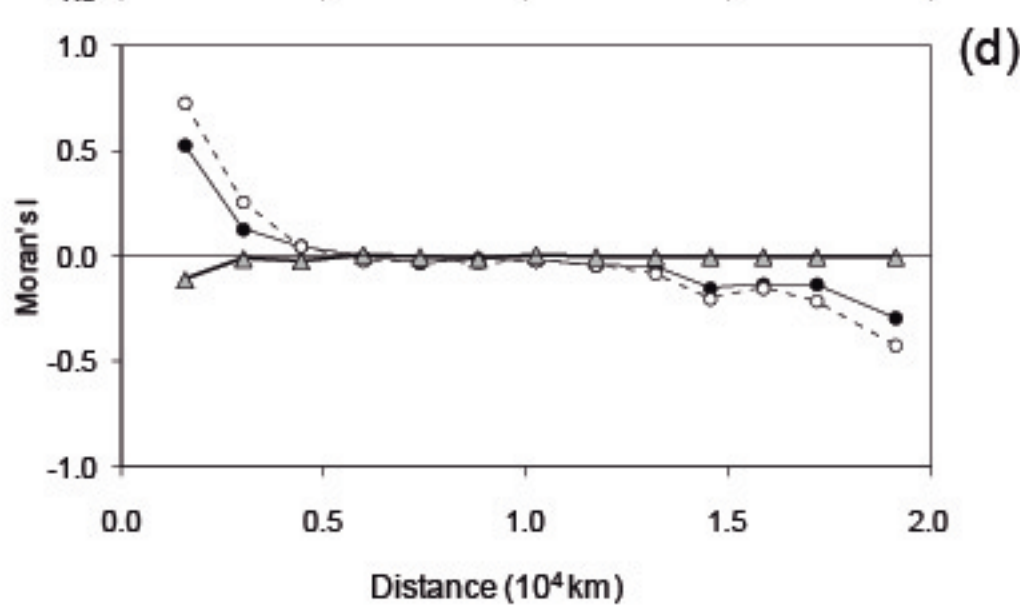
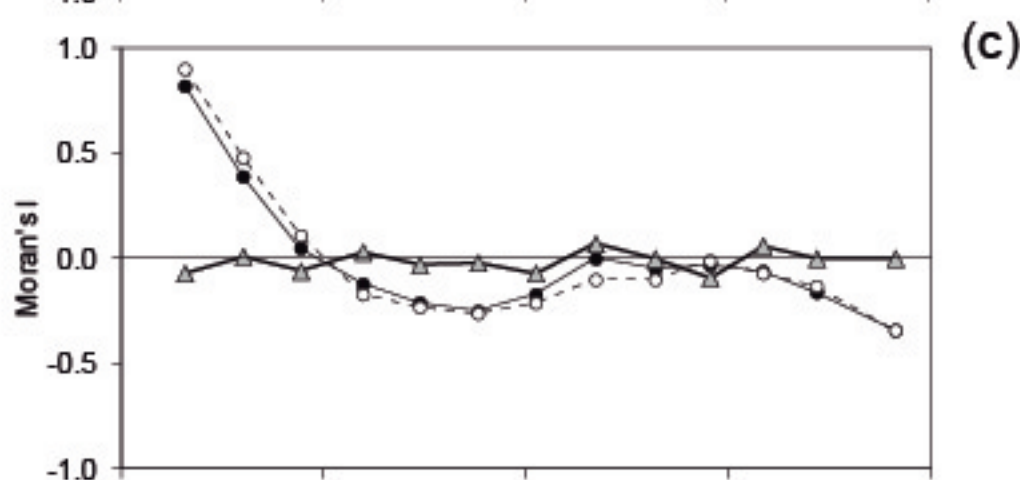
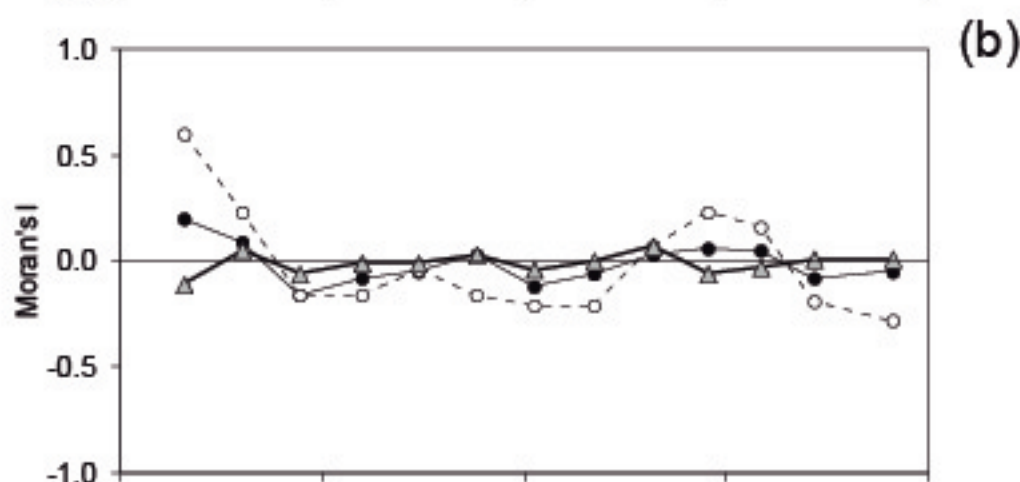
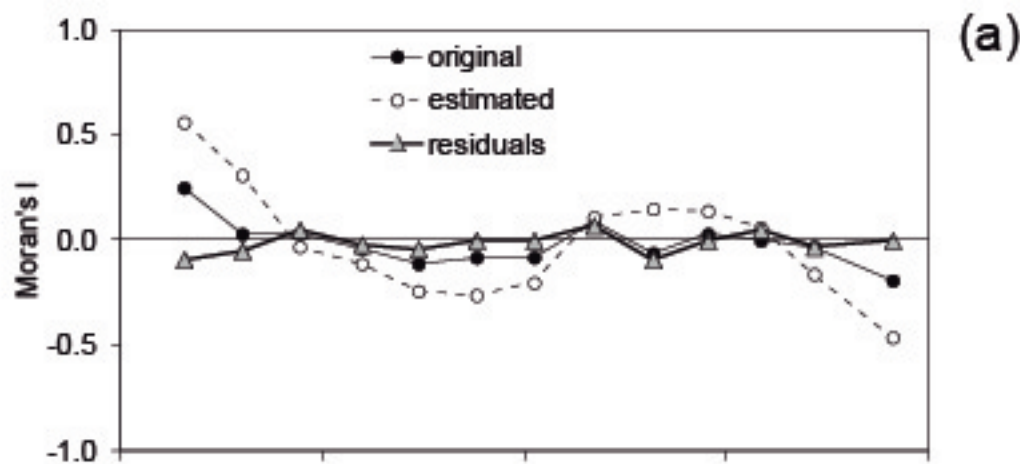




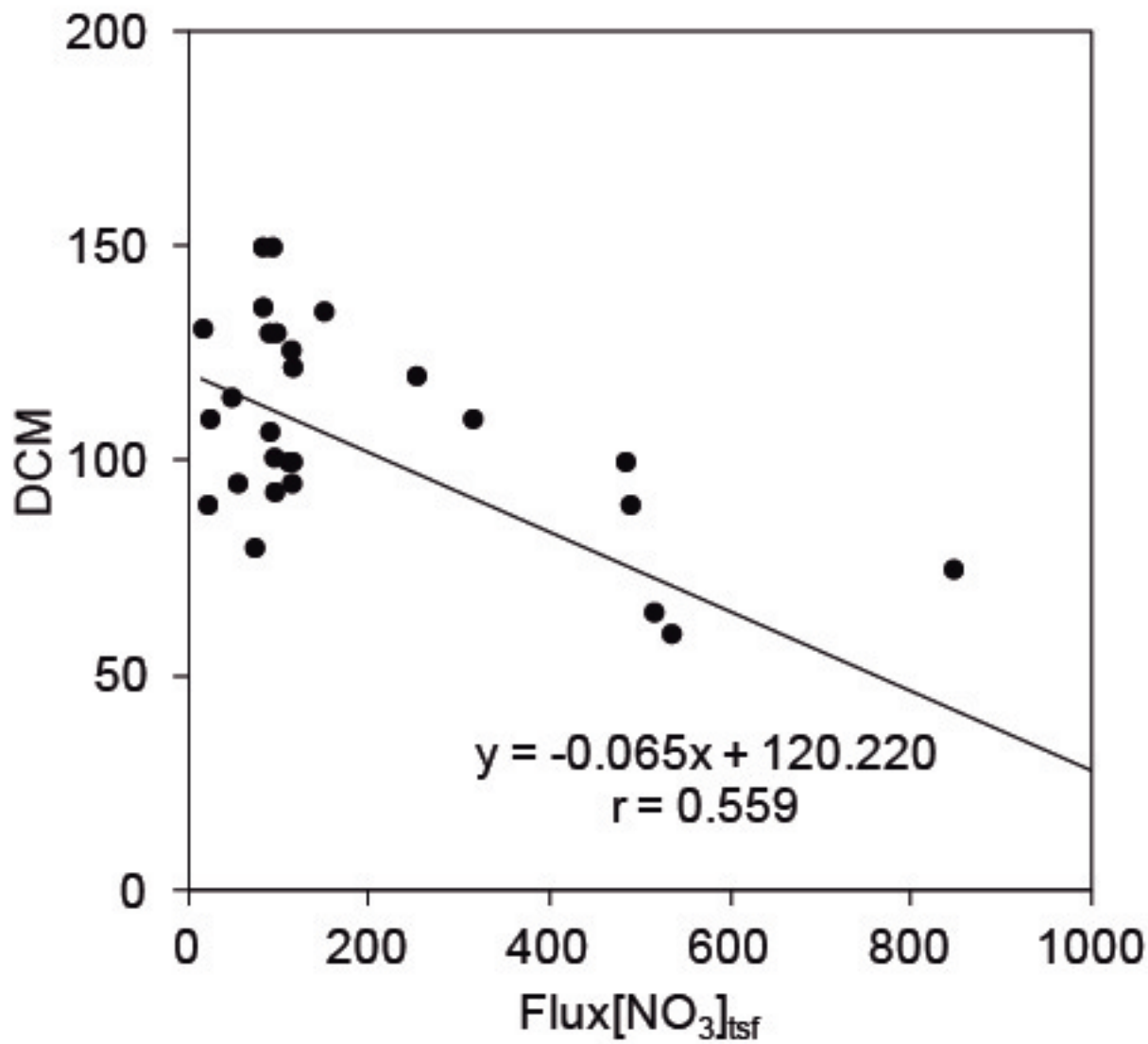






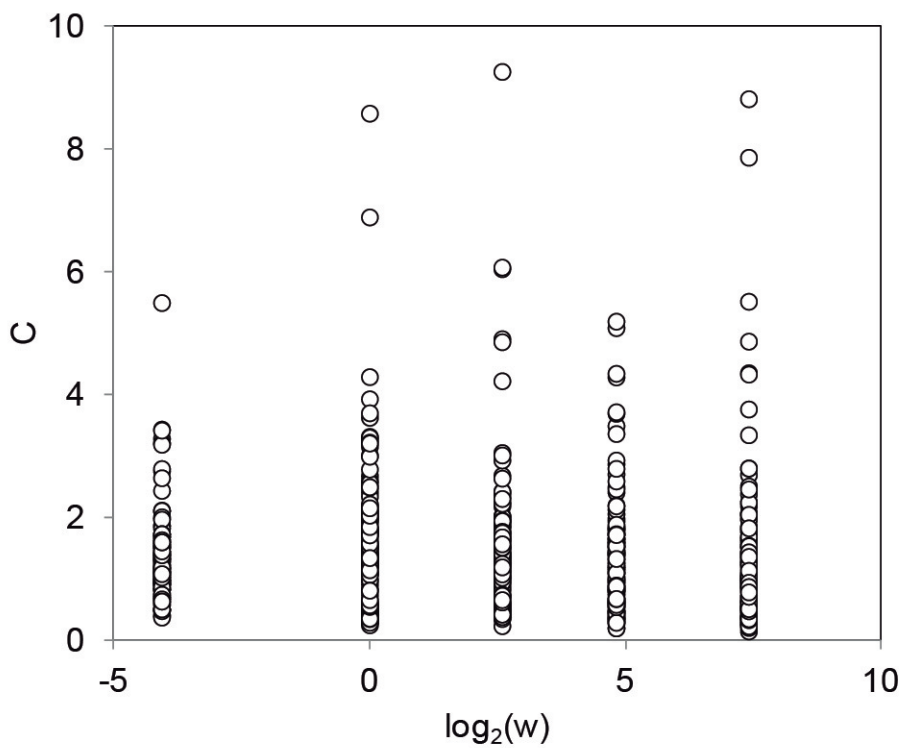








(a)



(b)

