Aluminum cycling in a tropical montane forest ecosystem in southern Ecuador

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14 Abstract

Growth limitation induced by Al toxicity is believed to commonly occur in tropical forests, 15 16 although a direct proof is frequently lacking. To test for the general assumption of Al toxicity, Al, 17 Ca, and Mg concentrations in precipitation, throughfall, stemflow, organic layer leachate, 18 mineral soil solutions, stream water, and the leaves of 17 native tree species were analyzed. We 19 calculated AI fluxes, analyzed temporal trends and modeled AI speciation in the litter leachate 20 and mineral soil solutions. We assessed potential Al toxicity based on soil base saturation, Al concentrations, Ca:Al and Mg:Al molar ratios and Al speciation in soil solution as well as Al 21 22 concentrations and Ca:Al and Mg:Al molar ratios in tree leaves. High Al fluxes in litterfall (8.77±1.3 to 14.2±1.9 kg ha⁻¹ yr⁻¹, mean±SE) indicate a high Al circulation through the ecosystem. 23 24 The fraction of exchangeable and potentially plant-available Al in mineral soils was high, being a 25 likely reason for a low root length density in the mineral soil. However, Al concentrations in all solutions were consistently below critical values and Ca:Al molar and the Ca²⁺:Al_{inorganic} molar 26 27 ratios in the organic layer leachate and soil solutions were above 1, the suggested threshold for

Al toxicity. Except for two Al-accumulating and one non-accumulating tree species, the Ca:Al molar ratios in tree leaves were above the Al toxicity threshold of 12.5. Our results demonstrate a high Al cycling through the vegetation partly because of the presence of some Al accumulator plants. However, there was little indication of an Al toxicity risk in soil and of acute Al toxicity in plants likely reflecting that tree species are well adapted to the environmental conditions at our study site and thus hardly prone to Al toxicity.

34 Keywords

35 tropical forest ecosystems, aluminum fluxes, aluminum toxicity, Al speciation, molar Ca:Al ratios

1 Introduction

Many plant species are sensitive to high concentrations of the phytotoxic Al³⁺, AlOH²⁺, or 37 38 AlOH₂⁺ and various other inorganic Al complexes which can occur in soil solutions at pH values <39 5.5 (Alleoni et al., 2010; Delhaize and Ryan, 1995; Kabata-Pendias and Pendias, 2001; Macdonald 40 and Martin, 1988). Aluminum phytotoxicity contributes to forest decline in temperate forests (Cronan, 1989; Farr et al., 2009; Godbold et al., 1988). In tropical montane forests, pH usually 41 42 ranges betwen 4 and 5 and Al toxicity was suggested to contribute to low biomass production and slow nutrient-cycling rates (Bruijnzeel, 2001; Bruijnzeel and Veneklaas, 1998; Hafkenscheid, 43 44 2000).

The Al fluxes in an ecosystem vary strongly depending on tree species (coniferous, deciduous), the climate conditions (temperate, tropical) and soil properties like texture, organic C concentrations and pH (Table 1). Aluminum inputs depend on dust deposition and amount of precipitation. Our literature review revealed that the highest Al fluxes with litterfall occur in tropical environments while the highest Al fluxes in soil solution were reported in acidified temperate forests, because of locally low soil pH (Table 1).

The pH is the most important control of Al concentrations in soil solution and acid deposition is a main driver of Al fluxes in a forest (Mulder, 1988). Thus, seasonal acid deposition originating from Amazonian forest fires (Boy et al., 2008a) and the increasing NH₄⁺ deposition with subsequent nitrification already in the forest canopy and nitrate leaching through the ecosystem resulted in acidification of the organic layer leachate at our study site (Wilcke et al., 56 2013), which will probably also couple back to Al fluxes in the system.

57 A hydroponic experiment with saplings of three different tree species typical for the south 58 Ecuadorian montane forests has shown that Al toxicity thresholds (EC10 values) are between 126 59 and 376 μ M Al in solution (Rehmus et al., 2014). However, knowledge of Al concentrations in soil 60 solutions alone is not sufficient to judge the threat of Al toxicity, because Al speciation is crucial 61 for toxic effects (Alleoni et al., 2010). If the solid Al pool is limited, which is particularly the case 62 in the organic layer compared to the mineral soil, ligand complexation in the solid and dissolved 63 phases leads to detoxification of AI^{3+} . Several studies demonstrated an alleviation of AI toxicity by 64 the formation of organo-Al complexes with dissolved organic matter in solution (Alleoni et al., 65 2010; Drabek et al., 2005; Hernandez-Soriano et al., 2013; Vieira et al., 2009) and in the solid 66 organic matter (Álvarez et al., 2012; Eimil-Fraga et al., 2015). At our study site, most nutrients are stored in the thick organic layers (Wilcke et al., 2002) where also 51 to 76 % of the fine root length 67 68 is located (Soethe et al., 2006). Previous studies showed that depending on dissolved organic 69 matter concentrations, 97 % to almost 100 % of the Al in organic layer solution is organically 70 bound in complexes and nontoxic (Wullaert et al., 2013).

71 A variety of indices based on chemical composition of the soil solid phase, soil solution, 72 and plant tissue can be used to estimate Al stress of an ecosystem (Álvarez et al., 2005). One 73 commonly used approach to estimate the threat of Al stress to plants is the Ca:Al molar ratio in 74 plant tissue and soil solution and the base saturation of the soil (Cronan and Grigal, 1995), 75 because the Ca-Al antagonism may disturb the Ca nutrition at high Al concentrations (Rengel, 76 1992). According to Cronan and Grigal (1995), indices for a 50 % risk of adverse impacts on tree growth induced by AI stress are a Ca²⁺:Al_{inorganic} (sum of inorganic AI species) molar ratio of \leq 1.0 77 78 in soil solution, a Ca:Al molar ratio of \leq 12.5 in the foliar tissue, and a soil base saturation \leq 15 % 79 of the effective cation-exchange capacity (ECEC). Aluminum can also affect the Mg nutrition of 80 plants (Kidd and Proctor, 2000; Kinraide, 2003). Reduced Mg concentrations in needles of Picea 81 abies (L.) H.Karst. in an in-situ experiment with elevated Al concentrations (up to 500 μ M) in soil 82 solution were revealed by De Wit et al. (2010). In a previous hydroponic experiment, impaired Mg translocation to the leaves and possibly reduced photosynthesis was suggested as a reason 83 84 for reduced shoot biomass production under Al stress (Rehmus et al., 2015).

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We analyzed Al fluxes and cycling in a tropical mountain forest in southern Ecuador and

tested soil, soil solution and plant leaves for indications of Al toxicity to answer the followingquestions:

- 881. Is Al cycling enhanced in tropical montane forests because of acid soils and89elevated litterfall production?
- 90 2. Do toxicity indicators point at negative effects of Al on plant growth?

91 2 Materials & Methods

92 **2.1** Study site and sampling procedures

The study site is an approx. 9 ha-large microcatchment (MC 2) between 1900 and 2010 m a.s.l. (above sea level) on 30–50° steep slopes on the north-facing part of the Rio San Francisco valley (Boy and Wilcke, 2008). The ecosystem flux measurements in the forest are concentrated along three ca. 20 m-long transects covering about 10 m in elevational difference at 1900 – 1910, 1950 – 1960 and 2000 – 2010 m a.s.l. (MC 2.1, MC 2.2, and MC 2.3, Figure 1). We considered the three measurement transects as replicates to account for the spatial variation in the study catchment.

100 Incident precipitation and throughfall were collected with Hellmann-type collectors on a 101 clearing and at the three transects in the forest, respectively. Stemflow was collected on five 102 representative large trees in MC 2.1 and surface flow was measured with a V-shaped weir at the 103 outlet of the stream draining MC 2. Further instrumentation (to collect litter leachate, soil 104 solutions, and litterfall) was placed along three measurement transects (MC 2.1 - 2.3). The 105 following equipment was established at each of the three replicate transects: Three litter 106 collectors with the dimension 0.3 m x 0.3 m and 0.5 mm mesh size, three zero-tension lysimeters 107 below the organic layer to collect organic layer leachate from below the Oi, Oe, and Oa horizons 108 (LL), and three suction cups at each of the 0.15 and 0.3 m depths of the mineral soil to collect 109 mineral soil solution (SS15 and SS30, respectively). Weekly sample replicates were bulked to a 110 composite sample per sample type (solution or litterfall) and measurement site prior to chemical 111 characterization. Solution samples were analyzed in weekly resolution while litterfall samples 112 were bulked to monthly samples before chemical analysis. Soil water content was measured with 113 FDR (frequency domain reflectometry) probes at transect MC 2.1 at the 0.1, 0.2, 0.3, and 0.4 m 114 depths. We calculated monthly element concentration means of incident precipitation, 115 throughfall precipitation, stemflow, and surface flow from April 1998 to March 2003, of soil 116 solutions from May 2000 to April 2003, and of the organic layer leachate from April 1998 to 117 December 2007 and September 2009 to April 2010. Our time series included smaller gaps (6 % of 118 time) because of missing samples. Samples of fresh tree leaves representing the most abundant 119 tree species in the highly biodiverse study area (Homeier et al., 2002) were collected in two 120 sampling campaigns between October 2005 and February 2006 and in October 2011 from 21 and 121 9 individual trees, respectively (17 tree species in total, Table 2). Young leaves were sampled 122 randomly from the tree crown. The base and Al saturation of the effective cation exchange 123 capacities of the mineral soils were determined as the sum of the charge equivalents of Ca, Mg, 124 Na and K or Al, respectively, divided by sum of the charge equivalents of Ca, K, Mg, Na, Mn and 125 Al extracted with 1 M NH₄NO₃ (soil:solution ratio 1:25) from the A horizons of 23 soils. Three of 126 the soil samples were collected at our measurement transects (each composited from three 127 sampling sites, Wilcke et al., 2001), 10 soils were located along an elevational transect near the 128 stream draining the watershed between 1880 and 2100 m a.s.l. (valley bottom) and further 10 129 soils were located along an elevational transect near the ridge between 1890 and 2110 m a.s.l. 130 (Wilcke et al., 2010). The latter two transects were considered to represent a large part of the soil 131 variation in the study catchment. More detailed information about the study site and sampling is 132 given in the Supporting Material (SM).

133 **2.2 Chemical analyses**

134 Within 24 h of collection of the precipitation, throughfall, stemflow, soil solution, and 135 stream water samples were analyzed first for electric conductivity (EC) (ProfiLine Cond 3110, 136 WTW GmbH, Weilheim, Germany) and then for pH (Sentix HWS, WTW GmbH, Weilheim, 137 Germany), filtered (ashless filters with pore size 4-7 µm, folded filter type 389; Munktell & Filtrak 138 GmbH, Bärenstein, Germany) and frozen at the day of sampling until analysis. To ensure that the 139 time gap between sampling and analysis had no effect on pH, selected samples were analyzed 140 immediately and after 24 and 48 hours and showed no change in pH (Table S1). The samples were 141 kept frozen until further chemical analyses. Aluminum concentrations were determined with 142 inductively-coupled plasma mass spectrometry (ICP-MS, VG PlasmaQuad PG2 Turbo Plus, Thermo

143 Fisher Scientific, Waltham, USA and 7700x Agilent Technologies, Frankfurt am Main, Germany). 144 The Ca, Mg, K, Na and Al concentrations in leaves and extracts of soils and suspended particulate 145 matter of stream water were analyzed with atomic absorption spectroscopy (AAS, SpectraAA400, 146 Varian, Darmstadt, Germany and Zeenit700P, Analytik Jena, Jena, Germany). Fluoride 147 concentrations were determined with an ion-sensitive electrode (WTW Inolab pH/Ion 735 with a 148 WTW F800 electrode, WTW, Weilheim, Germany) after addition of TISAB III (Fluka Analytical) by 149 standard addition. Concentrations of Cl⁻ were determined with a Cl⁻-specific ion electrode (Orion 150 9617 BN, Thermo Fisher Scientific, Waltham, USA) after adjustment of the ionic strength with a solution of 1.13 g L⁻¹ NaNO₃+2 mL L⁻¹ surfactant (Triton X-100, 50% solution) during the first 3 151 years on a segmented Continuous Flow Analyzer (CFA, AutoAnalyzer 3 HR, SEAL Analytical, 152 Germany). Concentrations of, NH₄⁺, NO₃⁻, and PO₄³⁻ were determined photometrically with a CFA 153 154 (initially a San plus, Skalar, Breda, Netherlands device was used and later the AutoAnalyzer 3 HR). 155 Sulfate was determined by ion chromatography (Dionex ICS-900, Thermo Scientific, Waltham, 156 MA, USA). Total organic carbon (TOC) concentrations were analyzed with a TOC-5050 (Shimadzu, 157 Düsseldorf, Germany).

158 The litter and leaf samples were digested in a closed vessel microwave system (MARS 159 Xpress, Kamp-Lintfort, Germany and MLS Ethos, Leutkirch, Germany, respectively) after drying 160 and homogenization with an agate ball mill. The litter and leaf samples of 2005/2006 were 161 digested with 65 % HNO₃ and leaf samples of 2011 with 69 % HNO₃/30 % H₂O₂/48 % HF and subsequently with 5 % H₃BO₃ to complex residual HF (16 HNO₃:6 H₂O₂:1 HF:10 H₃BO₃ v/v/v/v). 162 163 Concentrates of particulate matter in stream water were digested with 65 % HNO₃/48 % HF (4:1 164 v/v). The quality of digestions and measurements of Al, Ca and Mg concentrations was controlled 165 with the help of the certified reference material (CRM) BCR-100 (beech leaves IRMM, Geel, 166 Belgium). The accuracy of Ca and Al measurement in both types of digests was within ± 5% and 167 that of Mg within $\pm 10\%$ deviation of the certified value. The pH values of the solid soil samples 168 were determined in 0.01 M CaCl₂ (soil:solution ratio 1:2.5). The detection limits of our chemical 169 analyses are given in Table S2.

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171 2.3 Calculations and Al speciation modeling

172 The base saturation and saturation of exchangeable Al was calculated as the proportion 173 of charge equivalent of extractable Ca + K + Mg + Na and Al of the ECEC. The Al fluxes were 174 calculated for 5 consecutive hydrological years from April 1998 to March 2003 (in the case of 175 soil solutions 3 years from May 2000 to April 2003). More details about the calculation of fluxes 176 are given in the supporting material. Surface flow was modeled with the hydrological catchment 177 model TOPMODEL (Beven et al. 1995). Details about the parameterization and validation of the 178 model are given in Fleischbein et al. (2006). Soil water fluxes were calculated from climate data 179 and soil moisture measurements with a soil water budget model as described by Boy et al. 180 (2008b). Data gaps of soil water fluxes (because of lacking soil water contents) were substituted 181 with the help of a regression model of modeled weekly soil water fluxes on measured weekly throughfall volumes (R² = 0.85). The Al canopy budget and total Al deposition was estimated 182 183 with the canopy budget model of Ulrich (1983) assuming Cl⁻ as inert tracer (Boy and Wilcke, 184 2008) as described in the SM.

185 Potential drivers of Al concentrations are soil acidity and concentrations of complexing agents including TOC, F⁻, and SO₄⁻². Wullaert et al. (2013) showed that only TOC occurred in 186 187 relevant concentrations for Al complexation. Complexes of other ligands (Cl⁻, F⁻, NO₃⁻, PO₄³⁻, 188 SO₄²⁻) hardly contributed to total Al concentrations. Furthermore, total Al concentrations are 189 influenced by dilution/concentration effects which we addressed by including soil moisture 190 contents in our analysis. Multiple regressions were calculated among monthly Al concentrations 191 and Ca:Al molar ratios, respectively, with pH, TOC concentrations and soil moisture (only in the 192 organic layer leachate (see SM for more details).

193 The Al speciation was calculated using Visual MINTEQ (VMINTEQ, Version 3.0 beta, J.P. 194 Gustafsson) for LL (n=176), SS15 (n=41), and SS30 (n=50) samples, for which a complete data set 195 was available. The F⁻ concentrations were determined for a subset of 176 samples from all three 196 solution types and were in all cases below the limit of quantification (1.16 μ M) and in 89 % of the 197 cases even below the limit of detection (0.37 μ M). Thus, for the calculation we set F⁻ 198 concentrations of all samples to half of the detection limit (0.18 μ M). In the model, SO₄²⁻ 199 concentrations were set to 6.35 μ M (i.e. the mean sulfate concentration of 31 measured samples 200 of all three solution types). Ionic strength was estimated from EC according to Griffin and Jurinak (1973). Within VMINTEQ, the NICA-Donnan model was used to assess complexation of Al with
humic substances. Details of the NICA-Donnan model are given in Kinniburgh et al. (1996). NICADonnan properties of metal complexation by organic acids were taken from the literature (Milne
et al., 2003). An active DOM/DOC ratio of 2 was assumed and the dissolved organic acids were
adjusted to be 100 % fulvic acids (Tipping and Carter, 2011).

206 **3 Results**

207 **3.1 Al fluxes**

208 Mean annual AI fluxes at each of the three measurement transects were lowest in bulk, 209 dry, and throughfall deposition (< 0.6 kg ha⁻¹ yr⁻¹, Figure 1). Stemflow contributed negligibly to the 210 Al fluxes reaching the soil (< 0.05 - 5.4 %). Soil deposition by litterfall was 9 to 18 times higher 211 than that of the sum of all dissolved aboveground Al fluxes. The mean annual litter input at MC 2 212 from April 1998 to March 2003 was 10.0±0.3 t ha-1 yr-1. The mean annual canopy budget was 213 positive at each of the three measurement transects. The Al fluxes in organic layer leachate were 214 lower than in mineral soil solutions except at transect MC 2.3, where lowest Al fluxes occurred in 215 SS30. The multiple regression among Al concentrations (C_{AI}) and pH values, soil moisture content, and TOC concentrations (C_{TOC}) in organic layer leachate (Eq. 1) shows the highest partial 216 217 regression coefficient for the pH, highlighting the importance of pH for Al concentrations in soil 218 solutions (Figure S1).

219 220 Eq. 1 $C_{AI} = 7.88 - 0.88 \cdot pH - 0.03 \cdot soil moisture (vol.%) + 0.02 \cdot C_{TOC} (mgL^{-1})$ (R²= 0.31, p < 0.001, n = 75)

221 The net hydrological export of Al (i.e. SW-BD, Table 1) was slightly negative, indicating a net accumulation of Al, which was even more pronounced when the total catchment budget (i.e. 222 223 SW-[BD+DD]) was considered. Yet, the loss of Al as suspended particulate matter in stream 224 water averaged 28.8 kg ha⁻¹ yr⁻¹ (August 2000 – January 2003). From the sum of the net 225 hydrological export and the Al export in suspended particulate matter in stream water, the 226 weathering rate of Al can be estimated as 28.7 kg ha⁻¹ yr⁻¹ (Likens, 2013). This approach is based 227 on the assumption that the soil thickness is in steady state and that superficial erosion equals 228 soil formation at the subsoil-parent material border (Likens, 2013).

229 **3.2** Al toxicity indicators

The total Al concentrations in LL, SS15 and SS30 were usually higher in mineral soil solutions than in LL (Table 2). The total Al concentrations in precipitation, stemflow and stream water were below 1 μ M. The sum of the concentrations of inorganic Al species (Al_{inorg}) in LL, SS15 and SS30 ranged 0.00 - 30.9 μ M (Table S3). The fraction of organically bound Al was highest in LL (mean: 96 %) and decreased with soil depth (83 % in SS30).

Aluminum concentrations in leaves of the tree species in the study area were generally in the range of 0.05 – 0.14 (mg g⁻¹, Table 2). Two tree species, *Graffenrieda emarginata* (Ruiz & Pav.) Triana and *Miconia sp*, belonging to the family of Melastomataceae are Al accumulators (Jansen et al., 2002) and therefore had distinctly higher Al concentrations than the Al non-accumulating tree species.

The Ca²⁺:Al_{inorganic} and Mg²⁺:Al_{inorganic} molar ratios were distinctly higher in LL than in SS15 and SS30 solutions (Table S3). The Ca and Mg concentrations in the leaves ranged 1.7 - 12.5 and 1.2 - 5.3 (mg g⁻¹), respectively (Table 2). The Ca:Al molar ratios were consistently higher than 12.5 and ranged up to 144. Only the two Melastomataceae species and *Ocotea bentamiana* Mez. (Lauraceae) had Ca:Al molar ratios of 0.5 - 9.2, indicating potential Al toxicity.

The pH in soil A horizons ranged 3.2 – 5.2 (mean 3.7). The base saturation ranged 4.77 – 97.1 % (mean 41.4 %) and the Al saturation ranged 2.65 – 95.2 % (mean 57.7 %). Seven of the 23 soil A horizons had a base saturation of ECEC below 15 % and two soil A horizons had a base saturation between 15 and 16 % (Figure 2). Soils with low base saturation mainly occurred on the ridges (mean±SE base saturation on the ridges: 22.2±5.2 % and in the valley: 65.2±9.1 %) (Wilcke et al., 2010).

251 **4 Discussion**

252 **4.1 Al fluxes**

The Al fluxes with bulk and dry deposition were well in the range reported from other forests. They were higher than in the Hubbard Brook experimental forest (Likens, 2013) and a humid tropical ecosystem in Fiji (excluding cyclone events, Waterloo et al., 1997), similar to that of a temperate forest in the USA (Rustad and Cronan, 1995, Table 1) but low compared to other 257 tropical rain forests in Brazil (Mayer et al., 2000; Cornu et al., 1998) and a temperate forest in 258 Germany, which was affected by acid deposition (Matzner, 1989; Figure 1 and Table 1). The Al 259 bulk and dry deposition mostly derives from mineral dust, because Al is an ubiquitous element in 260 soils (Macdonald and Martin, 1988), but is also affected by the total amount of precipitation, 261 which is high compared to temperate but medium to low compared to other tropical forests. The 262 positive canopy budget indicates leaching of Al from the canopy. The leaching of cations is a result 263 of proton buffering in the canopy and release of cations to achieve electroneutrality (Matzner, 264 1989). The lowest Al fluxes occurred in stemflow, which were also lower than in other studies 265 (Figure 1). However, the input of Al and H⁺ ions by the stemflow to the soil is restricted to a small 266 area around the stem basis which might have a considerable impact on the local soil chemistry, 267 but minor importance for the total Al fluxes in the ecosystem (Koch and Matzner, 1993; Levia and 268 Frost, 2003).

269 The Al fluxes with litterfall were high compared to temperate forest ecosystems and most 270 study sites in the tropics (Table 1) and attributable to both, the overall high quantity of litterfall 271 and high Al concentrations in litterfall (Table 2 and Figure 1). Compared with a temperate forest 272 in Bavaria, Germany, mean annual litterfall was two and approximately five times that of a 273 deciduous and a coniferous forest, respectively, while Al concentrations in litterfall were around 274 10 times higher (Berg and Gerstberger, 2004). Litterfall and associated Al fluxes were lower at 275 some tropical montane forest (Jamaica, Table 1, Hafkenscheid, 2000) but similar to others 276 (Bolivia, Table 1; Gerold, 2008). The Al concentrations in litterfall were three times higher than 277 the Al concentrations in fresh foliage (Table 1), which is in agreement with Hafkenscheid (2000) 278 who reported enriched AI concentrations in litterfall compared to fresh leaves of the same species 279 in a tropical montane forest in Jamaica. They explained the Al enrichment as an attempt of the 280 trees to dispose excess AI via litterfall. Another reason is the withdrawal of mobile essential 281 nutrients prior to leaf fall, which will relatively enrich Al. However, another major explanation for 282 high Al concentrations in leaf litter is the wide distribution of Al-accumulating tree species. The 283 two families Melastomataceae and Rubiaceae, which are known to include Al-accumulating 284 species, belong to the most frequent families and *G. emarginata* is the most frequent tree species 285 in the studied forest (Homeier et al. 2002). However a guantification of the contribution of Al 286 accumulators to Al fluxes in litterfall would require samples in which leaf litter is differentiated according to the species which could not be realized in our study.

288 The AI fluxes with organic layer leachate and soil solutions had a wide range, covering 289 most of the span reported by other studies from similar soil depths (Figure 1 and Table 1). The Al 290 fluxes with the organic layer leachate and the soil solution are coupled to their pH (Eq. 1). In 291 stream water, the dissolved Al fluxes were low compared to two temperate forest ecosystems in 292 the USA (Likens, 2013; Rustad and Cronan, 1995), which is probably attributable to the high pH 293 of the stream water (Table 3), causing precipitation of Al e.g., as Al hydroxide. However, Al in 294 suspended particulate matter in stream water and thus also the AI weathering rate was 20 times 295 higher than e.g., that of the Hubbard Brook experimental forest (Likens 2013) in line with the 296 more pronounced tropical weathering regime at our study site than at Hubbard Brook.

297

4.2 Al toxicity indicators

298 The mean base saturation in the soil A horizons was low and the mean Al saturation high, 299 which implies reduced supply with base cations. The base saturation would even be lower if 300 exchangeable protons had been included to calculate the ECEC as it is done in part of the literature 301 (e.g., Cronan and Schofield, 1990). Because of the higher pH values in the valley bottom soils and 302 lateral addition of leached base metals from above-lying soils, the mean base saturation was 303 significantly higher in the valley bottom than the ridge top soils (Wilcke et al., 2010; Figure 2). The 304 ranges of base and Al saturation are complementary at respective pH values and show that the 305 fraction of exchangeable AI can be high (up to 95%). In mineral soils, AI might have an indirect 306 negative effect on plant nutrition via Al-P precipitation. However, the soil P status can be 307 improved by chelation of Al in organic complexes (Haynes and Mokolobate, 2001). The resulting 308 nutrient scarcity and high Al concentrations might be the reason for the low root length density 309 in the mineral soils and high root length density in the organic layers (Soethe et al., 2006). 310 Applying the critical base saturation of \leq 15 % proposed by Cronan and Grigal (1995), out of the 311 23 analyzed soils in our study approximately 30 % might pose a potential risk of Al toxicity, 312 particularly on the ridges, where 4 of the 7 soils with a BS < 15 % and two soils with marginally 313 higher BS (< 16 %) were located (Figure 2). The percentage of soils with critical base saturation 314 would even increase if exchangeable protons had been included in the ECEC.

The mean pH values in organic layer leachate and soil solutions were consistently < 5.5

316 and thus favorable for high Al concentrations (Table 3). The total Al concentrations which caused 317 a 10 % reduction in shoot biomass of saplings of three tree species from the same forest (EC10 318 values: 126 to 376 μ M Al, Rehmus et al., 2014) were not approached in any ecosystem solution. 319 To reach the EC10 values, an approximately 8 to 23-fold increase of the mean total Al 320 concentrations in organic layer leachate would be necessary. The lowest Al concentrations at 321 which a negative response of sensitive species was observed ranged from 35 to 170 μ M (Schaedle 322 et al., 1989). Thus the present total Al concentrations in the organic layer leachate can be 323 considered as nontoxic. In addition, the Al speciation modeling confirmed that up to 97 % of the 324 Al in the organic layer leachate is organically bound (Table S3). The remaining free Alinorg 325 concentration in the organic layer leachate of approximately 0.5 μM is unproblematic even for 326 Al-sensitive plant species (Wheeler et al., 1992). The mineral soil solutions had higher total Al and 327 lower DOC and thus highest concentrations of inorganic Al species. The mean Alinorganic 328 concentrations were above 0.5 µM in 30 cm depth at transect MC 2.1 and at both soil depths at 329 transects MC2.2 and 2.3 (Table S3) and thus Al toxicity might occur for sensitive plants. However, 330 high Alinorganic concentrations in in the mineral horizons will have a limited Al toxicity effect to local 331 tree species, because 51 to 76 % of the fine root length is concentrated in the organic layer 332 (Soethe et al., 2006). On the other hand, this distribution of root length density may be a reaction 333 to the higher Al concentrations in mineral soil solution compared to organic layer leachate.

Applying the Ca²⁺:Al_{inorganic} molar ratios in the organic layer leachate and the mineral soil 334 335 solutions as an indicator for a negative impact on the tree growth (Cronan and Grigal, 1995), only 336 few solutions might pose a 50 % and higher risk of Al stress (Table S3). Because Al has a higher 337 affinity for dissolved organic matter than Ca, the majority of the organic layer leachates and even the mineral soil solutions had a Ca:total Al and even Ca²⁺:Al_{inorganic} molar ratios far higher than the 338 suggested threshold of 1 (mean Ca²⁺:Al_{inorganic} molar ratios $1.17 - > 10^{36}$, Table S3), which renders 339 an Al effect on plant Ca uptake unlikely. Jorns and Hecht-Buchholz (1985) reported Mg deficiency 340 341 symptoms in Norway spruce (P. abies) at Mg:Al molar ratios of < 0.2. In the hydroponic 342 experiment with species from our study site, Al-induced toxicity symptoms occurred at > 300 μM 343 Al and Mg:total Al molar ratios in nutrient solution were < 0.2 at treatments \geq 600 μ M Al (Rehmus 344 et al., 2014). Thus, Mg deficiency might be a reason for reduced biomass production and Mg:Al 345 molar ratios might be a suitable indicator for Al stress. The Mg:Al molar ratios of SS15 and SS30 were < 0.2 but distinctly higher in all other ecosystem solutions, suggesting problematic Mg:Al
 molar ratios again only in the mineral soil (Table S2). However, calculating Mg²⁺:Al_{inorganic} molar
 ratios resulted in far higher values also in the mineral soil solutions (Table S3).

349 The Ca concentrations in tree leaves were all in the range required for optimal plant 350 growth $(1 - > 50 \text{ mg g}^{-1}; \text{ Marschner, 2012})$. The Mg concentrations in the leaves (Table 2) are 351 comparable to Mg concentrations in leaves of Brazilian Cerrado trees (0.71 – 2.1 mg g⁻¹; Lilienfein 352 et al., 2001) and in tree leaves of a tropical seasonal rain forest in southwest China (3.2 - 5.4 mg)353 g⁻¹; Shanmughavel et al., 2001). In leaves of 4 and 8 of the 17 analyzed tree species in our study, 354 Mg concentrations were below the limiting values required for optimal growth of crop plants reported by Amberger (1996) ($2 - 50 \text{ mg g}^{-1}$) and Marschner (2012) ($1.5 - 3.5 \text{ mg g}^{-1}$), respectively. 355 356 Thus, Mg deficiency might affect growth of some of the tree species. However, low Mg 357 concentrations in trees might primarily result from Mg-poor bedrock and only secondarily from 358 the Mg-Al antagonism. According to Cronan and Grigal (1995), the Ca:Al molar ratios of the two 359 Melastomataceae species and the Lauraceae species which were < 12.5 (Table 2) would mean a 360 50 % and higher risk of Al stress. However, beneficial effects of Al on plant growth in hydroponic 361 experiments with Al-accumulating tree species, adapted to acid soils, have been reported (Osaki 362 et al. 1997, Watanabe and Okada 2005). Only O. bentamiana, which is Al non-accumulating, might 363 suffer immediate Al stress.

364 **5 Conclusions**

365 We conclude that

Al fluxes in the ecosystem are mainly controled by high precipitation, biomass production,
 presence of Al-accumulating plant species and soil pH. Partly high Al fluxes, particularly in
 litterfall, illustrate a high Al circulation through the ecosystem possibly influenced by the
 abundance of Al-accumulating plants.

The fraction of exchangeable Al on cation exchanger surfaces and the concentrations of
 dissolved Al_{inorganic} were high in mineral soils. Together with the problematic Mg:Al ratios,
 Al toxicity is a likely reason for a low root length density in the mineral soil. The
 concentrations of inorganic Al species in the organic layer leachate were low and most
 probably below toxicity levels for trees and Ca:Al and Mg:Al molar ratios were

375 unproblematic. As plants have their main rooting zone in the organic layers, Al toxicity 376 might be of minor importance. The Al concentrations in tree leaves were high in Al-377 accumulating, but in a normal range for non-accumulating species. The Ca:Al molar ratios 378 in leaves of most Al non-accumulating tree species were unproblematic. Hence, the trees 379 of our study ecosystem seems to be well adapted to the given Al concentrations.

380

381 Supplementary Material

382 More information about the study site, sampling procedures, calculations, one figure and 3

tables can be found in the Supplementary Material.

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394 **Conflict of interest**

395 The authors declare no competing financial or personal interests.

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Table 1

Aluminum fluxes (kg ha⁻¹ yr⁻¹) of bulk (BD) and dry deposition (DD), throughfall (TF), stemflow (SF), litterfall (LF*), organic layer leachate (LL), mineral soil solution (SS**), stream water (SW), and suspended particulate matter loss with stream water (PL) from the literature. * total litterfall (t ha⁻¹ yr⁻¹) in parentheses, ** soil depth (m) in parentheses

		Al fluxes (kg ha ⁻¹ yr ⁻¹)								
Reference	Ecosystem	BD	DD	TF	SF	LF*	LL	SS**	SW	PL
Current study	tropical montane forest, Ecuador	0.2	0.2	0.5	0.01	11.4 (10.0)	5.0	10.19 (0.15), 6.02 (0.30)	0.18	28.8
Likens (2013)	northern hardwood forest, USA	< 0.01							2.79	1.38
Rustad and	northern red spruce	0.2		0.06	0.03	0.65	21		26	
Cronan (1995)	forest, USA	0.2		0.00	0.05	0.05	2.1		2.0	
Berg and	deciduous forest, Germany					0.98				
(2004)						(5.45)				
Matzner et al.	deciduous forest,							24(0.0)		
(2004)	Germany							2.4 (0.6)		
Berg and	coniferous forest									
Gerstberger	Germany					1.2 (2.14)				
(2004)	Germany									
Matzner et al.	coniferous forest,							17.2 – 26.9		
(2004)	Germany							(0.2 – 0.9)		
Matzner (1989)	temperate forest, Germany	1.2	0.9 – 1.3	1.6 – 2.9				17.6 – 52.7 (0.9)		
Cornu et al.	tropical lowland	1 4		0.6	0.02	2 62		27(04)		
(1998)	rainforest, Brazil	1.4		0.6	0.03	2.02		3.7 (0.4)		
Mayer et al.	rain forest, Brazil	50		3.2				26.5 - 43.5		
(2000)		5.2						(0.1 – 1)		
						1.50 -				
Hafkenscheid	tropical montane					5.22	1.9 –	16.1 – 35.1		
(2000)	forest, Jamaica					(6.47 –	4.9	(0.05 – 0.14)		
						6.16)				
Gerold (2008)	tropical montane forest, Bolivia					15.5				
Gerola (2000)						(12.2)				
						11.2 –				
Bergamini Scheer	Atlantic rain forest,					9.69				
et al. (2011)	Brazil					(6.37 –				
						3.01)				
Moraes et al.	Atlantic rain forest,					5.3 (6.31)				
(1999)	Brazil					5.0 (0.01)				

Table 2

Al, Ca and Mg concentrations and molar Ca:Al ratios in fresh leaves sampled between October 2005 and February 2006 and in October 2011 and in litterfall from April 1998 to March 2003. N is number of individual replicates, values represent mean ± standard error.

Species	Family	N	Ca	Mg	AI	Ca:Al molar ratio	
				(mg g ⁻¹)			
Sampling October 2	005 to February 200	6					
Purdiaea nutans Planch.	Cyrilliaceae	3	9.8 ± 0.8	2.6 ± 0.3	0.1 ± 0.0	53.7 ± 7.2	
Alchornea pearcei Britton.	Euphorbiaceae	3	6.6 ± 1.9	1.9 ± 0.3	0.1 ± 0.0	64.0 ± 16.5	
Graffenrieda emarginata (Ruiz & Pav.)	Melastomataceae	3	2.4 ± 0.2	1.6 ± 0.2	3.4 ± 0.4	0.5 ± 0.1	
Podocarpus oleifolius (Donex Lamb.)	Podocarpaceae	3	4.3 ± 0.7	1.4 ± 0.2	0.1 ± 0.0	37.1 ± 1.3	
Alazatea verticilata (Ruiz & Pav.)	Lythraceae	1	4.7	1.8	0.07	47.9	
Clusia ducoides (Engl.)	Clusiaceae	1	10.9	2.0	0.05	144	
Hyeronima moritziana (Mull. Arg.)	Euphorbiaceae	1	3.9	1.5	0.09	30.2	
<i>Ocotea aciphylla</i> (Nees) Mez.	Lauraceae	1	2.5	1.2	0.1	19.8	
Ocotea bentamiana Mez.	Lauraceae	1	1.7	1.2	0.12	9.2	
Miconia sp	Melastomataceae	1	7.7	3.3	1.64	3.2	
Elaeagia sp	Rubiaceae	1	3.7	1.6	0.10	25.4	
<i>Matayba inelegans</i> Spruce ex Radlk.	Sapindaceae	1	3.2	2.7	0.06	34.8	
Prunus opaca (Benth.) Walp.	Rosaceae	1	12.5	2.6	0.09	99.0	
Sampling C	October 2011						
Cedrela odorata L.	Meliaceae	1	6.5	2.3	0.10	42.9	
Cedrela sp	Meliaceae	2	9.1 ± 2.8	3.3 ± 0.3	0.12 ± 0.02	47.6 ± 9.5	
Heliocarpus americanus L.	Tiliaceae	3	7.6 ± 1.4	5.3 ± 0.7	0.14 ± 0.04	42.9 ± 13.3	
Tabebuia chrysantha (Jacq.) G. Nicholson	Bignoniaceae	3	5.8 ± 1.3	3.2 ± 0.4	0.07 ± 0.01	53.4 ± 11.5	
Litterfall	1998-2003						
litterfall		59	12.1 ± 0.2	4.0 ± 0.1	1.2 ± 0.1	8.5 ± 0.6	

Table 3

Range and mean of pH, mean (\pm SE) of Ca, Mg, Al, and TOC concentrations in bulk precipitation, throughfall, stemflow, organic layer leachate (LL), soil solutions in the 0.15 and 0.3 m soil depths (SS15 and SS30, respectively), and stream water at the three measurement transect in Microcatchment 2 (MC2.1 – 2.3) from April 1998 to March 2003 (LL April 1998 to April 2010) and in soil solutions from May 2000 to April 2003.

solution pH			Са	Mg	AI	тос
	range	mean		(μM)		(mg L ⁻¹)
precipitation MC 2	3.9 – 7.8	5.1	6.44 ± 0.62	4.44 ± 0.47	0.33 ± 0.04	5.28 ± 0.18
throughfall MC 2.1	4.5 – 8.2	6.2	25.1 ± 1.7	20.7 ± 1.3	1.39 ± 0.08	13.5 ± 0.4
throughfall MC 2.2	4.9 – 7.5	6.2	22.9 ± 1.2	20.4 ± 1.1	1.30 ± 0.08	14.3 ± 0.4
throughfall MC 2.3	4.2 - 8.0	6.0	41.9 ± 1.7	43.1 ± 2.0	1.47 ± 0.09	17.4 ± 0.8
stemflow MC 2.1	4.2 – 7.7	6.1	26.3 ± 1.4	22.0 ± 1.2	0.99 ± 0.07	18.3 ± 0.5
stream water MC 2	5.5 – 9.1	6.7	16.6 ± 0.7	18.2 ± 0.6	0.47 ± 0.07	4.48 ± 0.28
LL MC 2.1	3.8 - 6.6	4.4	31.9 ± 1.6	45.5 ± 1.9	27.18 ± 0.8	40.8 ± 0.9
LL MC 2.2	3.6 – 7.5	4.6	48.8 ± 2.5	58.2 ± 3.0	14.3 ± 0.9	33.0 ± 0.9
LL MC 2.3	3.3 - 6.9	5.0	111 ± 5.2	89.7 ± 4.8	4.42 ± 0.36	46.5 ± 1.9
SS15 MC 2.1	3.8 – 4.6	4.2	7.99 ± 1.20	7.14 ± 0.92	73.0 ± 3.02	27.0 ± 0.6
SS15 MC 2.2	4.0 - 5.6	4.4	6.19 ± 0.77	24.3 ± 1.9	12.4 ± 0.94	7.80 ± 0.23
SS15 MC 2.3	4.4 – 5.2	4.9	31.6 ± 1.9	16.6 ± 1.4	9.35 ± 0.32	10.7 ± 0.3
SS30 MC 2.1	4.1 - 4.9	4.4	3.53 ± 0.50	3.70 ± 0.55	36.8 ± 1.9	15.5 ± 0.4
SS30 MC 2.2	4.2 – 5.6	4.5	8.71 ± 1.08	24.7 ± 3.5	13.3 ± 1.4	7.77 ± 0.44
SS30 MC 2.3	4.4 – 5.6	5.2	40.6 ± 11.4	21.2 ± 5.0	4.86 ± 0.28	7.59 ± 0.36

Figure 1. Schematic illustration of Al fluxes (kg ha⁻¹ yr⁻¹) in bulk (BD) and dry deposition (DD), throughfall (TF), litterfall (LF), stemflow (SF), organic layer leachate (LL), soil solution in the 0.15 (SS15) and 0.3 (SS30) m soil depths, and stream water (SW), and the Al canopy (CB) and dissolved Al catchment budgets (WB) for an approximately 9-ha large water catchment under tropical montane rain forest in southern Ecuador. Shown are arithmetic means of annual values (\pm SE) from 1998 to 2003 (n = 5), in case of soil solutions from 2000 to 2003 (n = 3).

Figure 2. Frequency histogram of base saturation (% of ECEC) of 23 A horizons (a), of selected 10 A horizons in the valley bottom (b), and selected 10 A horizons on the ridge top of a 9-ha large catchment (MC 2).

Figure 1



Figure 2

