Anthropogenic footprint of climate change in the June 2013 northern India flood

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Abstract

During 13-17 June 2013, heavy rainfall occurred in the northern Indian state of Uttarakhand and led to one of the worst floods in history and massive landslides, resulting in more than 5,000 casualties and a huge loss of property. In this study, meteorological and climatic conditions leading up to this rainfall event in 2013 and similar cases were analyzed for the period of 1979-2012. Attribution analysis was performed to identify the natural and anthropogenic influences on the climate anomalies using the historical single-forcing experiments in the Coupled Model Intercomparison Project Phase 5 (CMIP5). In addition, regional modeling experiments were carried out to quantify the role of the long-term climate trends in affecting the rainfall magnitude of the June 2013 event. It was found that (a) northern India has experienced increasingly large rainfall in June since the late 1980s, (b) the increase in rainfall appears to be associated with a tendency in the upper troposphere towards amplified short waves, and (c) the phasing of such amplified short waves is tied with increased green-house gases (GHGs) and aerosols. In addition, a regional modeling diagnosis attributed 60-90% of rainfall amounts in the June 2013 event to post-1980 climate trends.

Keywords: Extreme events, climate and weather interactions, greenhouse gas (GHG) forcing, synoptic wave train, CMIP5, WRF model, cold air intrusion

1. Introduction

During 13-17 June 2013, heavy rainfall occurred in the northern Indian state of Uttarakhand, located on the windward side of the Himalayan ranges. The torrential rain together with rapid snowmelt led to extreme flooding and widespread landslides, causing thousands of deaths and a huge loss of property (Dubey et al. 2013). In addition to the devastation in Uttarakhand, this event also affected other parts of India including Himachal Pradesh, Haryana, Delhi and Uttar Pradesh, as well as western Nepal and parts of Tibet (Dubey et al. 2013). In recent years, similar heavy rainfall and widespread flood events have become increasingly frequent in northern South Asia. For example, an extreme rainfall event occurred in northern Pakistan during July 2010, resulting in floods that killed about 3,000 and affected around 20 million people (Hong et al. 2011; Lau and Kim 2012; Wang et al. 2011b). More recently (2-6 September 2014), some regions in India (Jammu and Kashmir) and Pakistan (Azad Kashmir, Gilgit-Baltistan and Punjab) underwent extreme floods caused by heavy rainfall, leading to more than 500 deaths (Najar and Masood, 2014).

A number of recent studies have investigated these heavy rainfall events, but most studies focused on either the synoptic or the mesoscale meteorological conditions of individual events (e.g., Hong et al. 2011; Houze et al. 2011; Joseph et al. 2014; Martius et al. 2013); few studies have analyzed the large-scale features and long-term climate

linkages. A recent study (Singh et al. 2014) conducted statistical analysis and concluded that the June 2013 rainstorm in northern India was at least a century-scale event, and the probability for such an event to occur has increased in the present climate compared to the preindustrial climate. However, knowledge regarding the mechanisms leading to the reported increased probability in extreme rainfall is lacking. Isolating the climate change impact on any individual storm or rainfall event is challenging, but such information is necessary for disaster planning and mitigation. Thus, the goals of this study are to identify common features in the meteorological conditions accompanying the June 2013 event and to investigate the mechanism through which climate change influences similar rainfall events, using observational data and climate model simulations. The data and modeling system used in this research are described in Section 2. The results are presented in Section 3. A summary and discussions are provided in Section 4.

2. Data and Methods

2.1 Data

To depict evolution of the heavy rainfall cases, observational rainfall was obtained from the 3-hourly Climate Prediction Center (CPC) Morphing technique precipitation (CMORPH) (Joyce et al. 2004) with the resolution of 0.25° × 0.25°. Since the CMORPH exists only after December 2002, monthly global precipitation data from NOAA's Precipitation Reconstruction over Land (PREC/L) (Chen et al. 2002) for the period of 1948–present was used to analyze the long-term climatology and trend. The PREC/L dataset is based on the gauge observations over 17,000 stations worldwide, and the resolution used in this study is 1.0° × 1.0°. For meteorological variables including wind,

temperature, relative humidity, and geopotential height, the NCEP/NCAR Reanalysis (Kalnay et al. 1996) for the period 1948–present was used.

To perform detection and attribution analyses, we used the fully coupled climate model simulations in the Coupled Model Intercomparison Project Phase 5 (CMIP5) (Taylor et al. 2012). To isolate the climate change signal, four sets of the CMIP5 Historical Single-Forcing Experiments were used: (a) one driven solely by natural forcing (e.g., solar cycle and volcano) (denoted as NAT), (b) one forced solely by greenhouse gases (denoted as GHG), (c) one driven solely by aerosols forcing (denoted as Aero), and (d) one driven with all natural and anthropogenic forcing sources (Taylor et al. 2012). A total of 10 coupled models were used in this study, and the details of these models are listed in Table 1.

2.2 Regional climate model experiments

Simulations of the June 2013 Indian rainfall event were carried out using the Weather Research and Forecasting (WRF) model version 3.5 (http://www.wrf-model.org/index.php). Initial and lateral boundary conditions were obtained from the NCEP-DOE Reanalysis2 (Kanamitsu et al. 2002), which is 6-hourly data with a resolution of 2.5° × 2.5°. WRF simulations were conducted for the period of 1-21 June 2013 and the first 11 days were treated as spin-up. The model land use was derived from the United States Geological Survey (USGS) 24-category global 30-second dataset. The spatial resolution was set to 30 km, and the simulations used 30 vertical layers up to 50 mb. The physics parameterizations included the SBU-YLin scheme for microphysics (Lin and Colle 2011), CAM schemes for radiation (Collins et al. 2006), MYNN level 2.5

98 TKE scheme for the Planetary Boundary Layer (PBL) processes (Nakanishi and Niino
99 2006), and five-layer soil thermal diffusion scheme for land surface processes.

100 To isolate the effects of climate change on the June 2013 Indian rainfall event,
101 two experiments were designed:

102 (1) *Control simulation* forced by the initial and boundary conditions (BC) from the
103 original NCEP-R2 data;

104 (2) *No-trend simulation* forced by the BC of the NCEP-R2 from which the post-1980
105 linear climate trends in all BC variables were removed. The assumption here is that
106 any long-term trend manifest in the troposphere contains signals that are traceable to
107 anthropogenic climate warming (which is supported by CMIP5 attribution analysis as
108 shown later). Although the long-term changes exhibited by different variables may
109 not be linearly correlated, we have shown in a previous study (Wang et al. 2011a) that
110 the nonlinear effect is generally negligible when it comes to this no-trend simulation
111 approach in South Asia.

112

113 **3. Results**

114 **3.1 The heavy rainfall event in June 2013**

115 To depict the large-scale environment associated with the June 2013 flood event,
116 we divided the evolution of the rainfall event (8-22 June 2013) into three periods: pre-
117 storm (8-12 June), storm (13-17), and post-storm (18-22) periods. In doing so, we
118 focused on the large-scale environment and its evolution. Figures 1a and b show the 5-
119 day averages of wind and vorticity fields during the pre-storm period at 200mb and
120 700mb, respectively. An upper-level ridge covered most of northern India (Figure 1a),

while the monsoon trough center (Figure 1b) was located on the western coast of the Indian peninsula and the Arabian Sea. Meanwhile a monsoon depression developed over the Bay of Bengal (BoB) as seen in the lower troposphere. These circulation patterns changed considerably during the storm period (Figures 1c and d): First, an upper-level tropospheric trough developed over northern India and appeared to be part of a short-wave train extending from the Mediterranean Sea to East Asia (Figure 1c). As indicated by Joseph et al. (2014), this trough over northern India induced cold air intrusion in the upper troposphere and subsequently enhanced instability in the region. In the lower troposphere, the BoB depression moved into the Indian subcontinent and merged with the monsoon trough, forming a strong cyclonic circulation over central and northern India. The northern branch of this cyclonic circulation apparently interacted with the Himalaya foothills, which provided orographic lifting and further enhanced rainfall in Uttarakhand and adjoining regions (Joseph et al. 2014). During the post-storm period (Figures 1e and f), the upper-level trough weakened and the lower-level cyclonic circulation over the Indian peninsula dissipated.

The aforementioned analyses show that the June 2013 extreme precipitation event was likely caused by several factors acting collaboratively: (a) deepening of the upper-level trough leading to increased baroclinicity, cold air intrusion aloft, and enhanced instability with warm and moist air beneath, (b) strong monsoon trough in the lower troposphere merged with a BoB monsoon depression, and (c) interaction of the circulation with a steep topography on the southern side of the Himalayan ranges. These regional meteorological conditions are symptomatically similar to those in other extreme rainfall events in northern South Asia that involved upper-level synoptic waves (Wang et

144 al. 2011b; Rasmussen et al. 2014).

145

146 **3.2 Comparison with events of similar circulation settings**

147 To investigate whether or not this June 2013 event is singular or recurrent in the
148 observational records and whether there is any systematic long-term change, we first
149 identified cases since 1979 that featured the upper-level circulation setting similar to that
150 of the June 2013 event. Since the midlatitude influence played a certain role (Joseph et al.
151 2014) and such an influence has appeared to intensify (Wang et al. 2011a), we designed
152 two selection criteria for the depiction of upper tropospheric circulations:

153 (1) For pattern recognition: The spatial correlation coefficient of 200 mb geopotential
154 height anomalies in the region (20°N-60°N, 0°E-150°E) between the June 2013 storm
155 period (13-17 June) and any given 5-day period is greater than 0.6 (i.e. with the p-
156 value < 0.001).

157 (2) For trough intensity: The area-averaged geopotential height at the center of the upper-
158 level trough (i.e., maximum vorticity in Fig. 1c to the northwest of Uttarakhand)
159 averaged over any given 5-day period is within 60-140% of that in the 13-17 June
160 2013 storm period.

161 These two criteria have to be met simultaneously to ensure proper identification of the
162 upper-tropospheric circulation pattern and trough strength that both resemble those in the
163 June 2013 event. Based on these criteria, only 5 cases were identified in the past 35 years
164 (1979-2013): 22-26 June 2004, 12-16 June 2007, 28 June-2 July 2009, 28 June-2 July
165 2010, and 28 June-2 July 2011. Apparently these cases only occurred in the last 10 years,
166 implying that this type of meteorological setting (or midlatitude influence) conducive to

167 extreme rainfall is likely influenced by climate change.

168 In Figure 2 we compared the CMORPH precipitation (Figures 2a and b) and
169 geopotential anomalies (Figures 2c-f) between the June 2013 event and the composite of
170 all 5 cases identified previously. The accumulated precipitation in the composite cases
171 does not show any significant amount in Uttarakhand (Figures 2b), even though the
172 upper-level short-wave train (Figure 2d) shares a similar pattern with the 2013 event
173 (Figure 2c). Why did these previous cases not produce rainfall as heavy as in June 2013
174 in Uttarakhand? An examination of the 700 mb geopotential height structure gives a hint
175 to this question: While the June 2013 event featured a strong monsoon trough (Figure 2e),
176 the composite cases are characterized by a weak monsoon trough across the Indian
177 subcontinent (Figure 2f). Altogether, these "similar but different" six cases reinforce the
178 previous claim that the June 2013 event occurred due to the unusual coupling of the
179 strong upper-level trough with a strong monsoon trough, and that these two anomalous
180 circulations at different levels do not always synchronize.

181 Figure 3a displays the time series of June precipitation averaged over Uttarakhand
182 (delineated with a box in Figure 1a), superimposed with an one-sided 20-year running
183 average (black line) and a linear trend after 1988 (red line). Apparently there has been an
184 increasing trend of precipitation during recent decades (with slope of 0.11 mm/day/year
185 at 99% statistical confidence). As a further examination, Figure 3b shows the spatial
186 pattern of the linear trend in the June 200 mb geopotential since 1988, reflecting the
187 maximum precipitation trend. Figure 3c shows the 200 mb geopotential anomalies during
188 the June 2013 event. A low pressure system is revealed in both Figure 3b and 3c to the
189 north of Pakistan, which facilitates upper-level cold air intrusion towards northern India

and western Nepal. This coincidence suggests that the upper-level short-wave train associated with the June 2013 event is embedded in a long-term change in upper tropospheric circulation structure. The coincidence also echoes the finding of Wang et al. (2011b), who analyzed the 2011 Pakistan flood in July and found that the post-1980 trend in the upper troposphere exhibited an amplified short-wave structure similar to that of the circulation anomalies during summer 2011. These observations are supportive of the emerging theory that the jet stream may have become increasingly "meandering"; this leads to an increase in extreme events worldwide (Francis and Vavrus 2012; Wang et al. 2013; Screen and Simmonds 2014).

3.3 Attribution of the climate trend

The next important question concerns the forcing mechanism that acts to strengthen the upper-level stationary waves near northern India. Here we analyzed the trend of the ensemble-mean 200 mb geopotential heights simulated by ten CMIP5 models for the period 1980-2005, and compared the results between the natural and GHG forcing experiments. As shown in Figure 4a, the simulation driven by all (natural and anthropogenic) forcing sources produced the 200 mb geopotential trends that are in reasonable agreement with the observation: i.e. an amplified wave train with an anomalous low center over central Asia and two anomalous high pressure centers located to the east and west. This result lends confidence in CMIP5 models' performance. However, Figure 4b shows that the simulation with only natural forcing produced a circulation structure that does not favor cold air intrusion over northern South Asia. In contrast, both GHG and aerosols forcing simulations (Figures 4c and 4d) produced the

213 200 mb geopotential trends that are in line with the observation with the deepened trough
214 to the north of Uttarakhand. This suggests that the increased greenhouse gases and likely
215 the increased aerosols collectively caused wave train pattern of the change in the upper-
216 level tropospheric flows. This result corresponds to the previous finding that all 5
217 previous cases having a similar upper-level circulation setting with the June 2013 event
218 occurred only in the last decade. Using three CMIP5 models, Wang et al. (2013) have
219 found that only the GHG forcing experiments produced the amplified short waves during
220 summer. In addition to the change in dynamics, we plotted in Figure 5 the June surface
221 (2-meter) temperature averaged over Uttarakhand superimposed with the post-1988 trend
222 (red line). The surface temperature in Uttarakhand only shows a mild warming trend that
223 did not pass the significance test ($p > 0.1$). This means that the upper tropospheric
224 cooling is relatively more important for the destabilization and associated precipitation
225 increase as revealed in Figures 3 and 4.

226 When it comes to attribution analysis, the mere use of observational data and
227 model free runs is not adequate to reach robust conclusions. As a complementary
228 approach, sensitivity experiments with WRF were performed (experimental design is
229 detailed in Section 2.2). Figures 6a-c show 5-day average precipitation during the storm
230 event (13-17 June) from the CMORPH as well as the control and no-trend experiments,
231 respectively. The control experiment (Figure 6b) produced rainfall in Uttarakhand that
232 agrees reasonably with the observation, while the no-trend experiment (Figure 6c) grossly
233 underestimated precipitation. Apparently, the removal of long-term trend in the WRF
234 boundary conditions considerably reduced the total storm rainfall. The ratio between the
235 no-trend and control experiments (Figure 6d) indicates a 60-90% reduction in rainfall

over Uttarakhand (boxed area), and such a reduction in rainfall amounts is considered attributable to the long-term climate change. As further attribution, Figure 6e shows the daily precipitation evolution averaged over Uttarakhand. While the control experiment produced a comparable amount of rainfall with the observation, albeit with a shifted timing (delay) by about one day, the no-trend experiment produced significantly reduced precipitation, i.e., less than 20% over the entire period of 13-17 June. This result illustrates that, although the rainfall event would still occur regardless of the climate trend or change, the post-1980 climate trend in the atmosphere has significantly aggravated the storm intensity.

The mechanism through which the climate trend has contributed to the severity of the June 2013 event is further illustrated through thermodynamic analysis. Figures 7a and b show the vertical profiles of WRF-simulated potential temperature lapse rate ($d\Theta/dp$) and relative humidity averaged over the Uttarakhand region during 13-17 June 2013. The $d\Theta/dp$ of no-trend experiment revealed a discernable stabilization in the 800-650 mb layer relative to the control experiment. The increase in stability in the no-trend experiment is compounded by the apparent drying below 700 mb amounting to ~10% in relative humidity (Figure 7b). Thus, the combination of stabilization and drying in the lower troposphere, in addition to the weakening of the upper-level trough and wave train, supports the substantial rainfall reduction simulated by the no-trend experiment due to reduced conditional instability.

4. Concluding remarks

We explored the meteorological and climatic conditions accompanying the June

2013 rainstorm event in northern India and analyzed past cases that feature similar upper-level circulation settings. The June 2013 event appears to be collaboratively generated by three factors: (a) an upper-level short-wave train with a cyclonic circulation over northern India leading to cold air intrusion, (b) a strong monsoon trough supplying moist air towards the Himalayan foothills, and (c) orographic lifting. The upper-level cold air intrusion enhances instability and subsequently increases rainfall intensity in the region. Furthermore, climate diagnoses suggest that the formation of the distinct short-wave train is not sporadic, but rather is reinforced by the long-term change in the upper troposphere. Based on the CMIP5 historical experiments, the upper-level wave train pattern revealed in the post-1980 trends is attributed to the increases in greenhouse gases and anthropogenic aerosols. Sensitivity experiments with the WRF model further indicated that the removal of the post-1980 trends in the forcing data leads to substantially reduced (~80%) precipitation in the flood region for the 5-day storm period. This estimated rainfall reduction is attributed to two prime factors: (1) suppressed cyclonic circulation in the upper troposphere restoring stability and (2) reduced moisture in the middle to lower troposphere. These processes favor the persistent increase in June rainfall over northern India after the mid-1980s and arguably contribute to the record amount of rainfall received in June 2013.

The conclusions reached in this study have implications for future flood management, water planning, and extreme weather prediction in northern South Asia. This study showed that as a result of anthropogenic climate change, the circulation structure has been modified in such a way that significantly aggravates rainstorm occurrences in northern South Asia, hence increasing the severity of floods. Also, the

occurrence of this June 2013 event during pre-monsoon season in northern South Asia, along with the circulation and precipitation trends in June, calls for prevention attention to increasingly frequent and strong rainstorms outside the core monsoon months (i.e., July-August). Adaptation measures such as developing strategies and policies for flood management in the face of climate-related extreme events are urged. In addition, the amplified upper-level stationary waves and associated dynamics as revealed in this study will need to be represented accurately in the forecasting tools.

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Figure 2: CMORPH precipitation averaged for (a) the storm event of 13-17 June 2013 and (b) the composite of five past events with similar circulation settings (see text). (c)-(d) Similar to (a)-(b) but for the 200 mb geopotential anomalies (HGT), with the long-term mean removed. (e)-(f) Same as (c)-(d) but for the 700 mb geopotential anomalies.

Figure 3: (a) Time series of June precipitation averaged over the Uttarakhand region (red box in Figure 1) superimposed with a 20-year running mean (black line) and a linear trend after 1988 (red line). The 2013 amount is highlighted in red, indicating its record status. (b) The spatial pattern of the post-1988 linear trend (slope) in the 200 mb geopotential height(HGT); unit is meter per 25 years. Stippling indicates regions exceeding 90% statistical confidence. (c) The 5-day mean 200 mb geopotential height anomalies of 13-17 June 2013.

Figure 4: The 1980-2005 linear trend in the 200 mb geopotential height (HGT) simulated by (a) the all forcing, (b) the natural forcing, (c) the GHG forcing, and (d) the aerosols forcing experiments of 10 CMIP5 models CMIP5 models. The unit is meter of total change over the 1980-2005 period. Stippling indicates regions exceeding 90% statistical confidence.

Figure 5: Time series of June surface temperature (2 meter) averaged over the Uttarakhand region (red box in Figure 1), superimposed with the post-1988 linear trend (red line).

Figure 6: Daily precipitation averaged for 13-17 June 2013 from (a) CMORPH, (b) the WRF control experiment, and (c) the no-trend experiment. (d) Percentage of precipitation reduction between the no-trend and control experiments; only the reduction in the no-trend experiment is shown. (e) 3-hour precipitation derived from CMORPH (blue), the control (black) and no-trend (red) experiments in Uttarakhand (boxed area).

Figure 7: Vertical profiles of (a) potential temperature lapse rate and (b) relative humidity averaged in Uttarakhand from the control (black) and no-trend (red) experiments averaged for 13-17 June 2013.

Table 1. CMIP5 (the Coupled Model Intercomparison Project Phase 5) models used in the attribution analysis

Acronym	Full name	Number of ensemble	Developers
CanESM	Canadian Centre for Climate modeling and Analysis The second Generation Earth System Model 2	5	Canadian Centre for Climate Modelling and Analysis
CCSM4	Community Climate System Model version 4	3	National Center for Atmospheric Research
CNRM-CM5	National Centre for Meteorological Research Coupled Model 5	6	Centre National de Recherches Meteorologiques /Centre Europeen de Recherche et Formation Avancees en Calcul Scientifique, France
GFDL-CM3	Geophysical Fluid Dynamics Laboratory Coupled Physical Model 3	3	NOAA, Geophysical Fluid Dynamics Laboratory
GFDL-ESM2	Geophysical Fluid Dynamics Laboratory Earth System Model 2	1	NOAA, Geophysical Fluid Dynamics Laboratory
CSIRO	Commonwealth Scientific and Industrial Research Organization	4	Commonwealth Scientific and Industrial Research Organization/Queensland Climate Change Centre of Excellence (CSIRO-QCCCE)
FGOALS	Flexible Global Ocean-Atmosphere-Land System	1	Institute of Atmospheric Physics, Chinese Academy of Sciences
GISS-E2	Goddard Institute for Space Studies Model E2	3	NASA, Goddard Institute for Space Studies
IPSL-CM5	Institute Pierre Simon Laplace Coupled Model 5	3	Institute Pierre-Simon Laplace
NorESM1	Norwegian Earth System Model 1	1	Norwegian Climate Centre (NCC)

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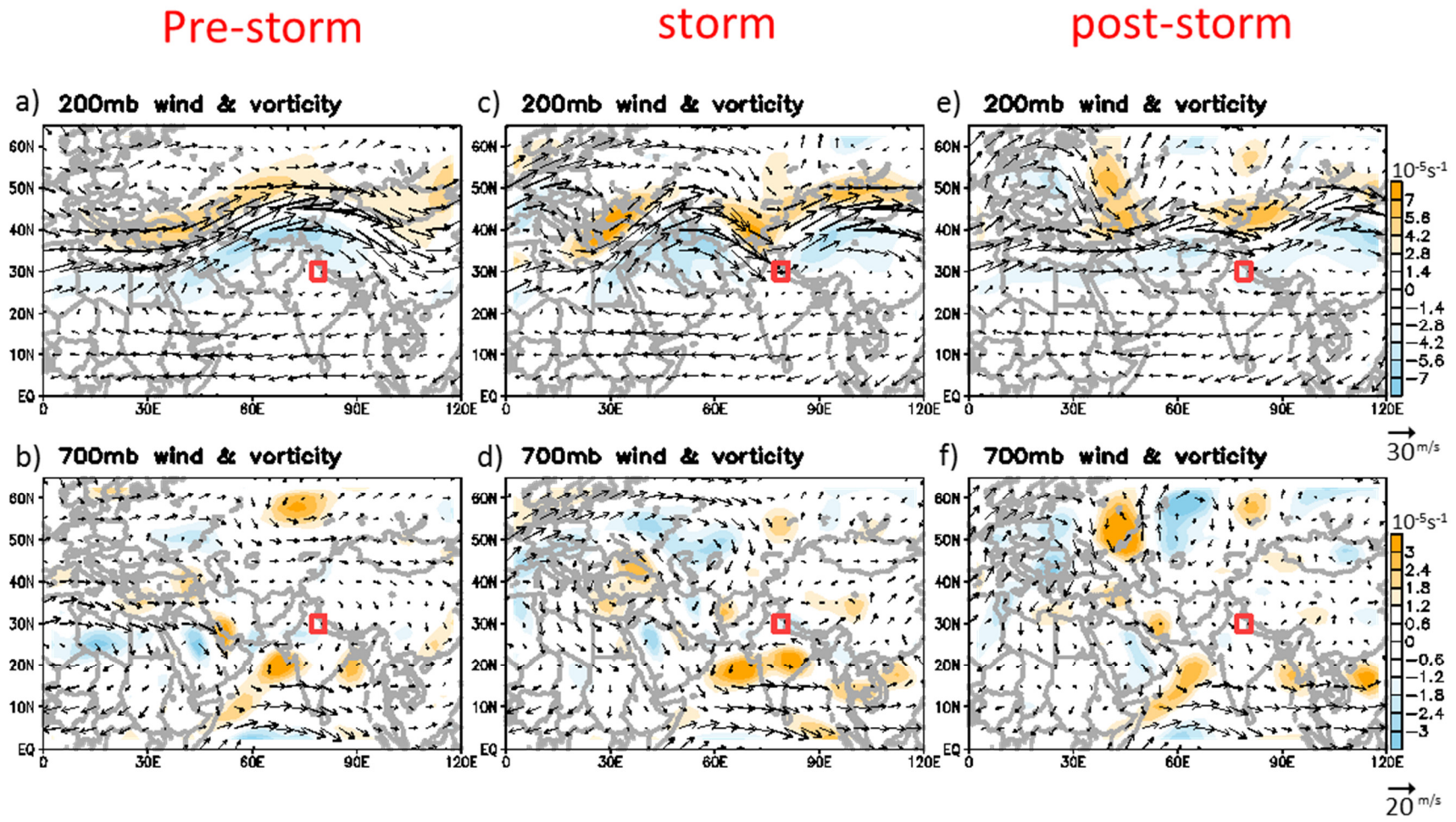

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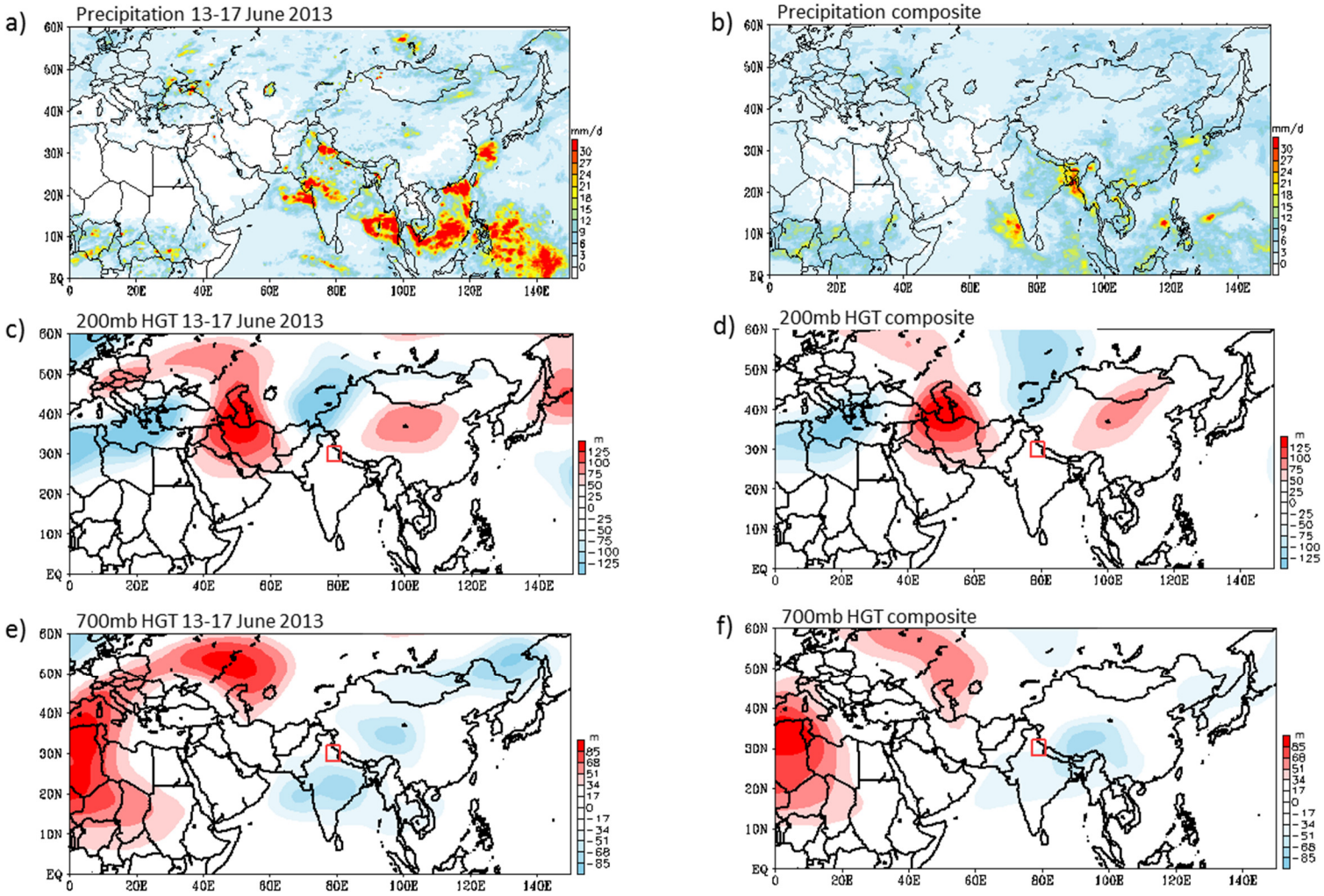

Figure 2: CMORPH precipitation averaged for (a) the storm event of 13-17 June 2013 and (b) the composite of five past events with similar circulation settings (see text). (c)-(d) Similar to (a)-(b) but for the 200 mb geopotential anomalies (HGT), with the long-term mean removed. (e)-(f) Same as (c)-(d) but for the 700 mb geopotential anomalies.

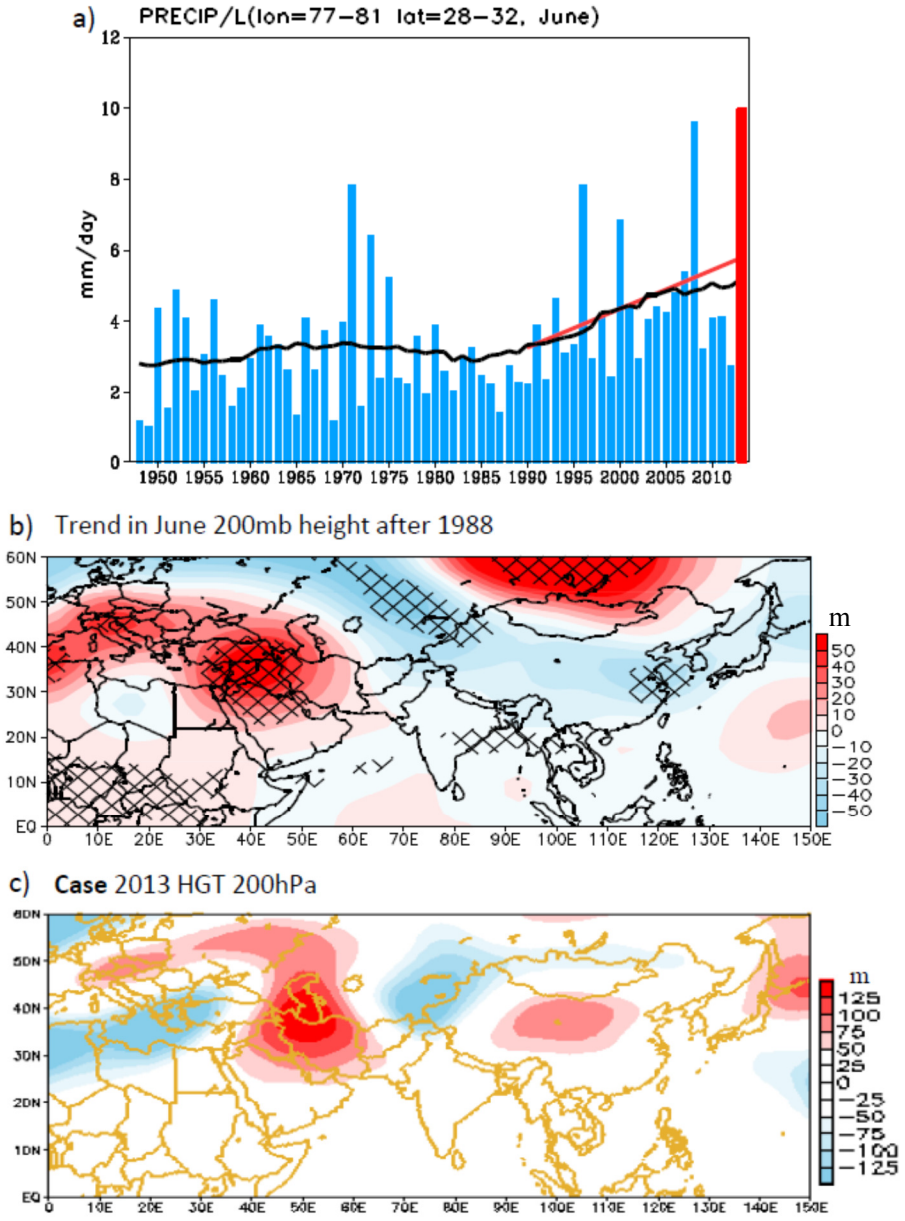

Figure 3: (a) Time series of June precipitation averaged over the Uttarakhand region (red box in Figure 1) superimposed with a 20-year running mean (black line) and a linear trend after 1988 (red line). The 2013 amount is highlighted in red, indicating its record status. (b) The spatial pattern of the post-1988 linear trend (slope) in the 200 mb geopotential height (HGT); unit is meter per 25 years. Stippling indicates regions exceeding 90% statistical confidence. (c) The 5-day mean 200 mb geopotential height anomalies of 13-17 June 2013.

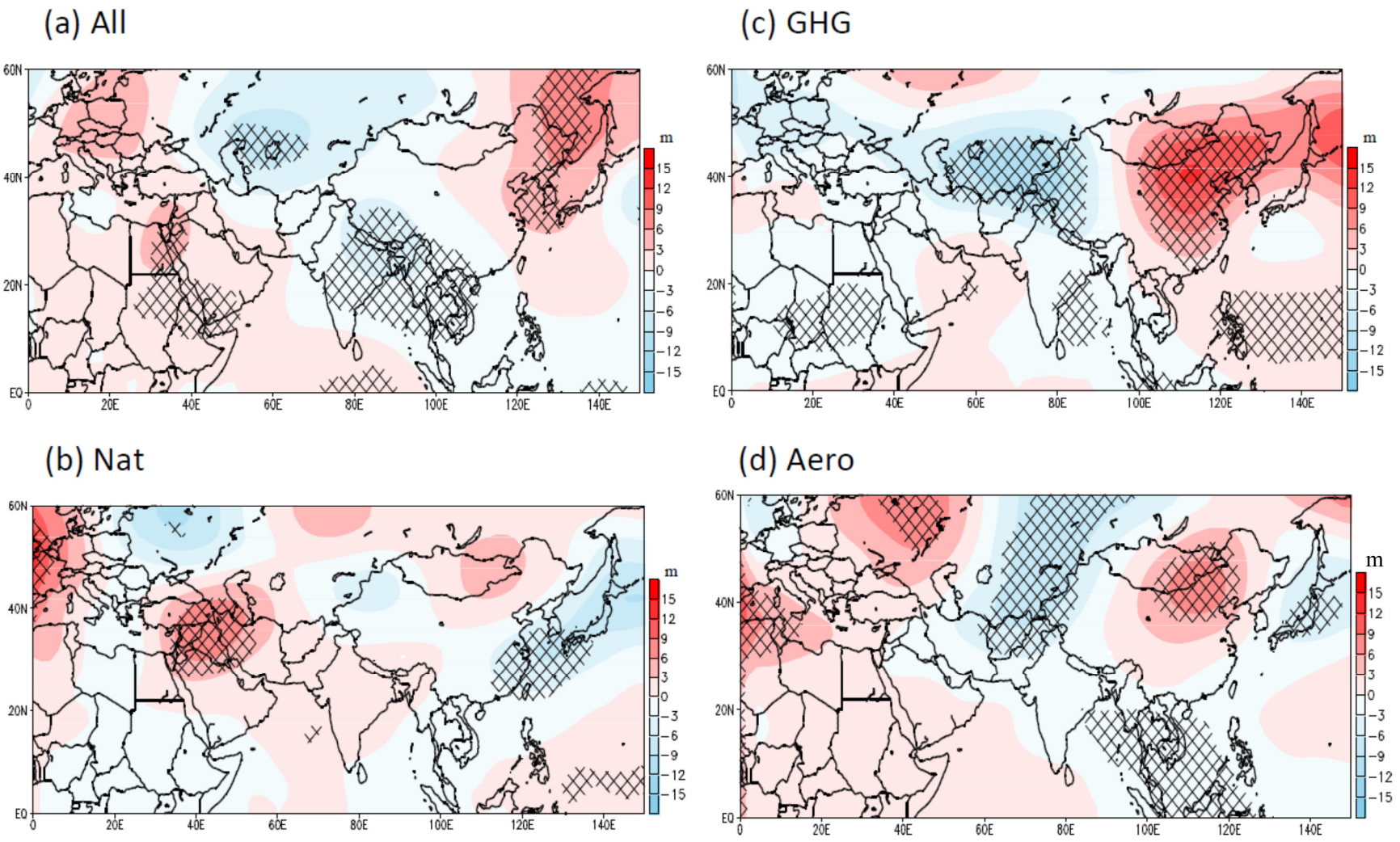

Figure 4: The 1980-2005 linear trend in the 200 mb geopotential height (HGT) simulated by (a) the all forcing, (b) the natural forcing, (c) the GHG forcing, and (d) the aerosols forcing experiments of 10 CMIP5 models CMIP5 models. The unit is meter of total change over the 1980-2005 period. Stippling indicates regions exceeding 90% statistical confidence.

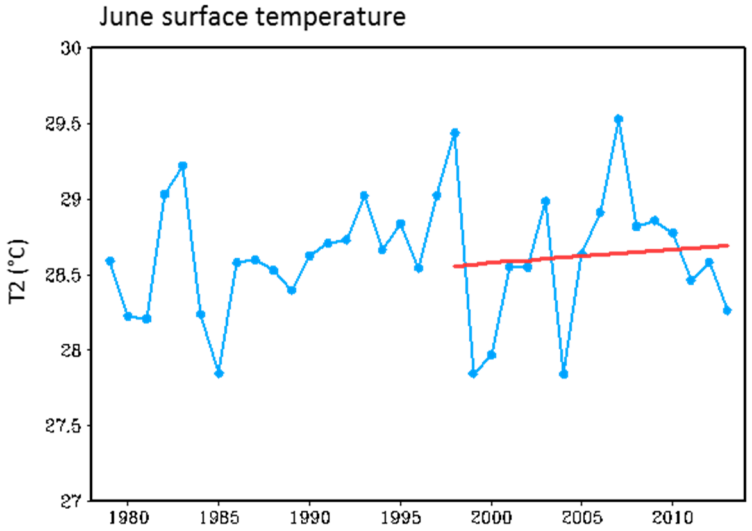

Figure 5: Time series of June surface temperature (2 meter) averaged over the Uttarakhand region (red box in Figure 1), superimposed with the post-1988 linear trend (red line).

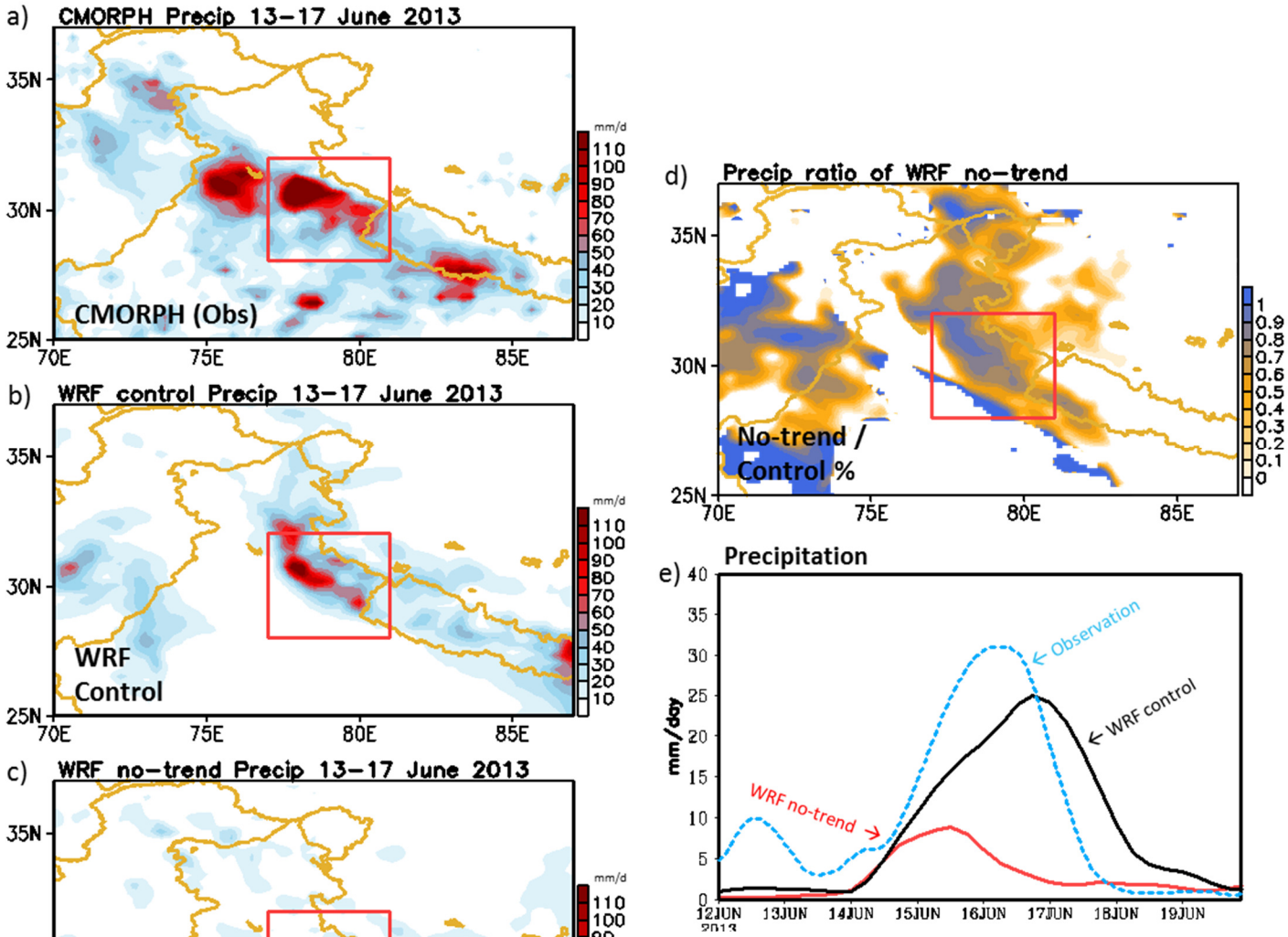

Figure 6: Daily precipitation averaged for 13-17 June 2013 from (a) CMORPH, (b) the WRF control experiment, and (c) the no-trend experiment. (d) Percentage of precipitation reduction between the no-trend and control experiments; only the reduction in the no-trend experiment is shown. (e) 3-hour precipitation derived from CMORPH (blue), the control (black) and no-trend (red) experiments in Uttarakhand (boxed area).

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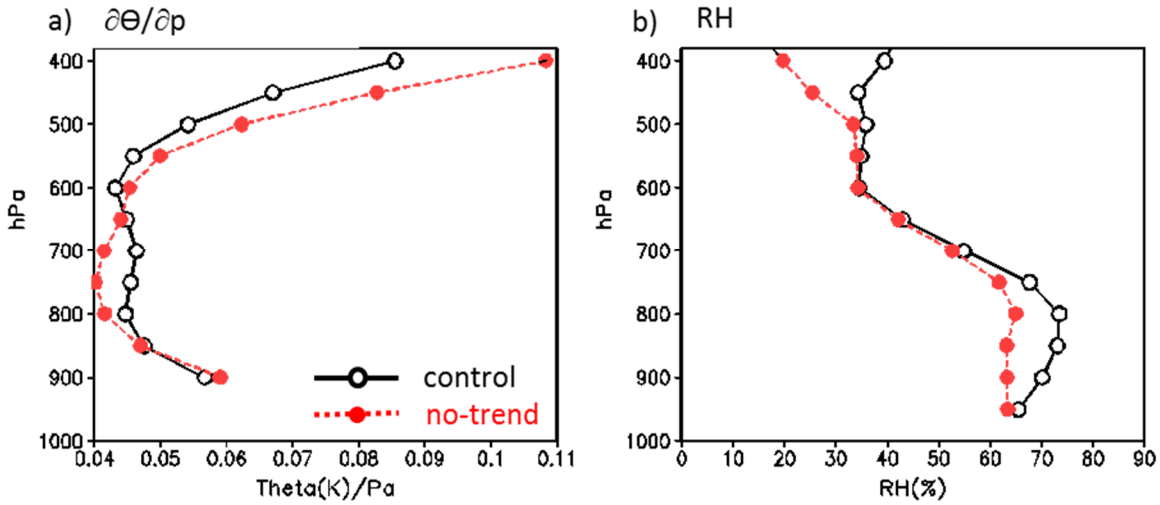

Figure 7: Vertical profiles of (a) potential temperature lapse rate and (b) relative humidity averaged in Uttarakhand from the control (black) and no-trend (red) experiments averaged for 13-17 June 2013.