

1 **Testing soil fertility of Prehispanic terraces at Viejo Sangayaico in**
2 **the upper Ica catchment of south-central highland Peru**

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22 **Abstract**

23 This study presents a pilot geoarchaeological investigation of terraced agricultural systems near San
24 Francisco de Sangayaico, in the upper Ica catchment of the Southern Peruvian Andes. It aims to assess the
25 evidence for soil fertility associated with agricultural strategies practiced throughout the Prehispanic,
26 Spanish colonial and modern occupations in this region. A series of twenty-two test pits were hand
27 excavated through two terraced field systems, and sampled to examine the changes in soil physical and
28 chemical characteristics down-profile and downslope.

29 This study provides the first geoarchaeological analyses of the agrarian soil system surrounding Viejo
30 Sangayaico in the upper Ica catchment. Results demonstrate that the soil system was much modified prior
31 to the creation of the terrace systems, probably about 900 years ago. This system was characterised by a
32 weakly acidic to slightly calcareous pH, a consistent but low electrical conductivity, reasonable-but-
33 variable phosphorus content, and a loamy soil texture with a component of weathered volcanic tonalite
34 parent material. The shallow terrace soil build-up on the slopes investigated indicates that slope
35 modification was as minimal as possible. Moreover, the relatively low frequencies of organic material and
36 phosphorus suggest that the terraces were not heavily fertilised in the past, making the stability and
37 management of the nutrient-rich topsoil vital.

38 The results of these excavations and soil fertility analyses are situated within the context of the wider
39 Andean ethno-historic and the archaeological record to address questions regarding how the terraces were
40 built and maintained over time. Agricultural terraces undoubtedly mitigated the effects of slope erosion
41 associated with cultivation. But, the terrace soil features observed at Sangayaico do not appear to be the
42 same as those documented in other geoarchaeological studies of Andean terrace systems. These contrasts
43 may be accounted for by a combination of differing geological substrate and hydrological conditions, as
44 well as variable trajectories in past soil development, erosion factors, manuring/field management
45 practices and crop selection.

46 Keywords: argillic tethering, geo-chemistry, micromorphology, soil fertility, terraces

47 **Introduction**

48 In this study the soil fertility of two Andean terrace systems and their associated buried soils
49 adjacent to the archaeological site of Viejo Sangayaico (or SAN1) is assessed (Fig. 1). It is set
50 within the Quebrada Marcaccarancca of the highland Olaya tributary of the Río Ica in the south-
51 central Andes of Peru. The site and its environs including ancillary settlements cover *c.* 30 km²
52 situated between 2800-4200 m.a.s.l., and were ethnically part of the Late Prehispanic Chocorvos
53 ‘nation’ (Rowe, 1946). This site is dated to Late Intermediate, Inca and Spanish colonial periods
54 with a range of radiocarbon dates from cal AD 1122 ± 81 (952+/-27 BP (OxA-30914-6) to cal
55 AD 1527 ± 88 (362+/-27 BP; OxA-30930-1). It is located at the transition between the low-lying
56 agricultural *quechua* ecozone (or irrigated, cultivated terrace zone at 2,300-3,500 m.a.s.l.) and
57 the increasingly agro-pastoralist *suní* (or dry field agricultural zone at 3,500-3,800 m.a.s.l. and
58 *puna* ecozones (or upland pasture zone at 3,800-4,800 m.a.s.l.) (Covey, 2006; D'Altroy, 2003;
59 Pulgar Vidal, 1946). The area exhibits significant technological investment in terracing and
60 irrigation canals for agriculture on the mountain slopes *quechua* zone, and high altitude check
61 dams in the high *puna* zone geared towards the creation of good pasturage. Viejo Sangayaico
62 appears to have been situated so as to control access to and from these ecozones and across the to
63 the northern Pisco Valley, an important hub of Chocorvos (AD 1000-1450) and later Inca (AD
64 1450-1532) and Spanish (post-AD 1532) occupations (Bueno Mendoza, 2003; Chauca and Lane,
65 2015; Huaman Oros and Lane, 2014; Lane et al., 2015). Although no direct dating evidence was
66 retrieved from the palaeosols and terrace systems at Sangayaico, it is highly likely that the
67 construction and use of the extensive bench terracing found in the area below *c.* 3500 m.a.s.l.
68 relates to the same periods as the adjacent occupations recorded and radiometrically dated.

69 Situated in an area rich in agricultural terracing, both abandoned and in use, this paper presents a
70 preliminary assessment of soil fertility of part of the terrace system around the Sangayaico sites,
71 as well as providing important comparative information concerning the creation and maintenance
72 of terraces in the Andean highlands.

73 Geologically, the area is composed primarily of tonalite parent material originating from
74 volcanic activity that occurred during the Cerozoic-Quaternary transition (Palacios Moncayo,
75 1994). Tonalite is primarily composed of andesine, biotite, hornblende, quartz, and orthoclase
76 minerals (Nettleton et al., 1970). The weathering of biotite in tonalite produces clay particles
77 such as kaolinite and vermiculite, as well as siliceous mica, and hornblende and quartz
78 weathering produces sand, silt and clay particles (ibid., 1970). It is these weathered mineral
79 constituents that define the substrate characteristics of the landscape surrounding Sangayaico.
80 Stone and coarse-fine sand-size fragments of this tonalite parent material are found throughout
81 the soil materials composing the terrace systems. The regional soils developed on this parent
82 rock range from luvisols with evidence of clay migration to weakly developed cambisols, with
83 leptosols on the steeper, rockier slopes (Gardi et al., 2015; WRB, 2014).

84 Soil fertility analyses are a crucial tool in the characterisation of past and present agro-
85 ecosystems (Sandor et al., 2007). Given the scale of human impact on the Andes (Denevan,
86 1992; Lentz, 2000), it is surprising that so few geoarchaeological studies of terrace systems exist
87 in this geographical area, especially given the large amount of work that has been done on
88 terraces more generally in Peru (i.e. Branch et al. 2007; de la Torre and Burga, 1986; Farrington,
89 1980; Gelles, 2000; Kemp et al. 2006; Kendall and Chepstow-Lusty, 2006; Kosok, 1965;
90 Mitchell and Guillet, 1994; Schjellerup, 1986; Trawick, 2003; Treacy, 1994; Williams, 2006).
91 This study goes some way to rectifying this gap in our knowledge and provides comparative soil

92 fertility analyses across several agricultural systems, thus providing important insights on the
93 interpretation of past agricultural strategies and other socio-cultural practices in the region
94 (Goodman-Elgar, 2009; Sandor et al., 2007; Wells, 2006).

95 The many previous studies of terraces elsewhere in the Americas have suggested that there is
96 firm evidence for terracing in the Pre-classic Mayan period *c.* 2000 years ago, reaching a peak in
97 the Late Classic period about 1300 years ago (Beach et al. 2015; Bonavia, 1967-1968; Denevan,
98 2001; Donkin, 1979). Terraces fulfil four main functions: providing a platform for a deep soil
99 matrix which facilitates cultivation, control of erosion, the creation of a sustainable micro-
100 climate and control of humidity (Treacy, 1994), as well as enhanced soil moisture and organic
101 matter contents. The classic image of Andean terracing and those found in the study area is that
102 of bench terraces, variously known as *andenes*, *patasi*, *bancales* and *takhanes* (Denevan 2001).
103 Around Sangayaico, the terrace walls are constructed using dry-walling known locally as *pirca*
104 masonry (Fig. 3). Although few of these terraces have been excavated, those that have such as in
105 the Paca, Cuzco, Colca and Chicha-Soras valleys were similarly built with either a single or
106 double course of *pirca* dry-stone masonry walls (Bonavia, 1967-1968; Brooks, 1998a; Denevan,
107 2001; Donkin, 1979; Goodman-Elgar, 2002; Kemp et al., 2006; Londoño, 2008; Schjellerup,
108 1986; Treacy, 1994).

109 Previous ge archaeological studies have observed a number of important characteristics of
110 Andean terrace systems. These include the burying of original agricultural horizons during
111 terrace construction, the preferential construction of terraces above argillic horizons - referred to
112 as tethering (Homburg et al., 2005), the application of fertiliser and the use of seasonal burning
113 (Table 1). In particular, the tethering of agricultural terraces to argillic horizons is known from
114 the Colca Canyon, and the Paca and Chicha-Soras valleys in Peru (Goodman-Elgar, 2008; Kemp

115 et al., 2006; Sandor and Eash, 1991, 1995). Their creation and maintenance results in the
116 formation of an anthropogenic topsoil covering the original agricultural horizons which have
117 been profoundly affected by physical, hydrological and geochemical alterations (Bryant and
118 Galbraith 2010). This may result in elevated phosphorous and organic matter levels, a decreased
119 pH and increased concentrations of illuvial silt, clay, and organic matter (Goodman-Elgar, 2008;
120 Sandor and Eash, 1991, 1995). Andean studies by most of these same authors also note the
121 presence of archaeological artefacts throughout the terrace sequence, thus suggesting the
122 repeated application of household waste as fertiliser through middening. There may also have
123 been corralling of animals to specific fields, most likely llamas or alpacas, and/or the collection
124 and direct application of manure from separate corrals. These techniques would have
125 supplemented the organic matter status of the field, increasing the nutrient content as well as the
126 water holding capacity of the soil, thereby enhancing the productivity of the landscape (Sandor
127 and Eash, 1991, 1995). In addition, there may be spikes of immobile phosphorous (Holliday and
128 Gartner, 2007; Sandor and Eash, 1995) and decreases in pH (to weakly acidic conditions) that
129 were counteracted by seasonal burning leading to the increased availability of phosphorus (P),
130 potassium (K) and carbonates. Regular burning would have also helped to remove deleterious
131 micro-flora and micro-fauna whilst depositing nutrient rich ash (Thomaz et al., 2014).

132 According to local Sangayaiqueño farmers, once the growing season was over animals grazed on
133 either the remains of the harvest or on the specially grown alfalfa (*Medicago sativa*), and then
134 the terrace fields were regularly burnt off. Growing seasons are short and are followed by long
135 fallow seasons of one to five years. These practices allow Sangayaiqueño farmers to maintain a
136 productive landscape without the use of artificial fertilisers.

137 At Sangayaico, two associated terrace systems were examined by targeted test pit excavations to
138 investigate changes in soil physical and chemical characteristics, both down-profile and
139 downslope. A total of 22 test pits were hand excavated across these terrace systems (Figs. 2-5),
140 described and sampled for physical, chemical and micromorphological analyses.

141 The results of this pilot study are then compared to other investigations of terrace systems in the
142 wider literature in order to address three main questions:

- 143 1. Is there evidence for buried agricultural horizons?
- 144 2. What past agricultural strategies (i.e. tethering to argillic horizon, fertiliser use, etc.)
145 are evident in the terrace systems associated to the Sangayaico site?
- 146 3. How have past and present agricultural strategies affected soil fertility?

147 **Survey and laboratory methodology**

148 In the landscape around Viejo Sangayaico two terrace systems (Fig. 2, A/red and B/blue) were
149 selected on the basis of their proximity to the SAN1 site and the apparent preservation of the
150 terrace architecture. Terrace system A was situated immediately downslope of SAN1 and
151 covered roughly 6 km², ranging from an elevation of 3585-3625 m.a.s.l., and had been left fallow
152 for the past three to five years. Below Terrace system A, Terrace system B was in use at the time
153 of excavation which limited excavation to three test pits. This system covers roughly 12 km²,
154 ranging in elevation from 3532-3577 m.a.s.l. Nineteen test pits were excavated (either 50 cm or
155 100 cm square) in the upper slope area (Terrace System A: TP1-4 & 6-20) along three parallel
156 transects perpendicular to the slope to the top of weathered tonalite bedrock, and a further three
157 test pits were excavated in a lower terrace system in a single transect (Terrace System B: TP5, 21
158 & 22) (Figs. 2-5). The test pit soil profiles were described (Table 2) and 84 bulk soil samples

159 were collected for physical and chemical analyses, and seven soil blocks taken for soil
160 micromorphological analysis.

161 The test pits revealed a quite consistent set of profiles through both terrace systems. The soil
162 profiles comprised either just thin ploughsoils over the weathered tonalite bedrock, or the
163 ploughsoil over a depleted zone of terrace made-ground which buried a remnant of a probable
164 former thin cambisol soil profile developed on the weathered tonalite bedrock. Small bulk and
165 micromorphological samples were collected from one 1x1 m test pit and of two 50x50 cm test
166 pits at six levels from the main representative soil horizons across the terrace system. The test
167 pits were dug to the base of the buried B horizon and/or top of the weathered tonalite parent
168 material (Figs. 4 & 5).

169 Samples were subjected to light grinding using a pestle and mortar before being sieved through a
170 2 mm mesh sieve, then processed using the suite of analyses described in Table 3 (with
171 references therein; Soil Survey Staff 1993), which includes, pH, electrical conductivity (EC),
172 organic matter content (loss-on-ignition), soil moisture content, particle size analysis,
173 phosphorus (P) content and soil micromorphology. Given the tonalitic parent material and
174 various sand-sized aggregates present in the soil samples, some methodological adaptations were
175 necessary to produce useful data. For the organic matter content and particle size analyses,
176 sodium hexametaphosphate was used as a deflocculant to achieve a pH of 9 and disperse the
177 sample fabric and break-down the clay (kaolinite) component (Devesa-Rey et al., 2011; Dwomo
178 and Dedzoe, 2010; El-Swaify, 1980; Gee and Bauder, 1986; Goodman-Elgar, 2008; Silva et al.,
179 2015), as well as a vortex and rotor mixer and sub-sampling from the mid-point of the
180 suspension for particle size analysis using the Malvern Mastersizer S Laser Diffraction Analyzer
181 (Gee and Bauder, 1986). To account for the inclusion of medium-grained sand in the samples

182 processed by laser diffraction, medium-sand values from wet-sieving and laser diffraction were
183 combined to decrease distribution errors.

184 Unfortunately without the use of radiocarbon or optically stimulated luminescence (OSL) dating,
185 the excavated terrace systems can only be relatively dated by assuming relationships to the
186 occupation of the associated Sangayaico site. There was no evident organic material from the
187 soil profiles in the test pits that was suitable for radiometric dating, and there were insufficiently
188 clear contacts between the terrace make-up deposits and possible buried soils to justify sampling
189 for OSL dating. Consequently the field systems only have a relative chronology through the
190 terrace systems association with the settlement sites on the Sangayaico ridge above. There is now
191 a series of radiocarbon dates ranging from cal AD 1100-1500 (OxA-30914/15/16, OxA-
192 30930/31) for the Late Intermediate site from the associated excavations at Sangayaico (Lane et
193 al., 2015). Another limitation met in the field was the fact that the entirety of the arable
194 landscape has been cultivated, leaving no natural controls to test against.

195 Finally, a series of seven blocks were taken from Test Pits 1, 2, 4 and 5 for soil
196 micromorphological analysis (Courty et al., 1989; Bullock et al., 1983; Murphy, 1986; Stoops,
197 2003, 2010). These aimed to be representative of, and characterise, the main stratigraphic
198 horizons present in the terrace system. Their analysis would serve to ground-truth the other
199 physical analyses, and indicate the pedogenic processes at work. This was part of a wider
200 geoarchaeological study of the upper Ica valley (French, 2015, pp. 54-62).

201 **Results**

202 The results of our study are summarised in Tables 2-7. The research questions considered are
203 addressed at the scale of the individual profile, transect, and terrace system. The quantitative
204 results demonstrate some down-profile and downslope trends (Table 6), and in combination with

205 the micromorphological analyses (Table 7), the assembled data create a clear picture of the
206 inherent soil characteristics and processes of the Sangayaico terrace soils.

207 The test pit profiles

208 For the majority of test pit profiles four soil horizons were evident (Table 2; Fig. 5). These
209 comprised a modern plough zone (or Ap) and an eluvial (or Eb) horizon, both fine sandy loams
210 which have developed in the upcast soil of variable thicknesses used to construct the terrace,
211 overlying a variable thickness (*c.* 8-25cm) of a buried, fine sandy/silty clay loam soil which is
212 probably a former A horizon (labelled bA2), all developed on the weathered tonalite bedrock. In
213 seven of the Test Pits (TP 3, 8, 9, 16, 18, 21 and 22), including most of Terrace System B
214 downslope (TP 21 and 22), there was no evidence of a buried soil or former A horizon surviving
215 at the base of the profile, with the terrace deposits situated directly on the bedrock. The test pits
216 in the lower Terrace System B also exhibited the greatest profile depths of *c.* 56-67cm.

217 pH and electrical conductivity

218 Soil pH ranged from weakly acidic to weakly calcareous with a range of values from 5.33 to 7.19
219 and an average of 6.34 (Tables 4-6). Terrace System A was more neutral in pH range; Terrace
220 System B was more weakly calcareous to neutral in range. Lower values of pH and concomitant
221 low concentrations of inorganic carbon are common among soil systems on volcanic substrates
222 such a tonalite (Nettleton et al., 1970). Down-profile, pH values varied little and mainly
223 remained in the neutral range, whereas down-slope the trend is for the profiles to become slightly
224 more calcareous.

225 Electrical conductivity ranged from 26.6 to 178.8 $\mu\text{S}/\text{m}$, averaging 57.1 $\mu\text{S}/\text{m}$ (Tables 4 and 5).
226 These values are all relatively low and do not suggest a high potential for elemental changes and
227 reactions taking place in this soil system. The highest EC values were from the uppermost

228 growing horizon, a common feature in arid environments due to the deposition of salts in the
229 topsoil during evapo-transpiration and the breakdown of organic matter (Meurisse et al., 1990;
230 Rhoades et al., 1999; Smith et al., 1996).

231 Soil moisture and organic contents

232 The soil moisture content of the air-dried soil ranged from 0.86 to 6.05%, averaging 2.39%, but
233 increased down-profile by as much as 4% (Tables 4-6). In Terrace System A, the soil moisture
234 content for the upper half of the profile was quite low but even, with the lower parts of the
235 profiles showing greater variability, fluctuating between 1.41 to 4.48% and 1.65 to 6.05%.

236 The soil organic matter data from the loss-on-ignition analysis (LOI 500 and Leco TruSpec)
237 ranged from 2.1 to 7.17%, with two outlier high values of 8.8% in the A horizons of TP 1 and
238 10.89% in TP18 (Tables 4-6). In Terrace System A, the soil organic matter content generally
239 remained quite stable to slightly decreasing by 1-4% down-profile, with slight enhancement in
240 the A horizon. In Terrace System B, the organic matter content remained quite steady with a
241 range of 4.79-5.73%, except for 8-10.54% high values in the Ah/Eb horizons of TP5.

242 Particle size analysis

243 Due to the variability of the sand- and silt-sized particle distribution, no clear patterns were
244 evident down-profile in either terrace system. However, the percentage distribution of the clay-
245 sized fraction generally decreased down-profile, with variability increasing downslope (only in
246 Terrace System A). In Terrace System A Row 1, there was a down-profile increase in the
247 distribution of clay, most notably between layers 3 and 4 or in the buried B soil, and between
248 Rows 3 and 1 the clay component increased to 11% with a coincident increase in silt content.

249 Terrace System B showed little soil textural variation downslope, with the middle terrace (Row
250 2) showing a higher distribution of sand at the expense of the silt- and clay-sized fractions.

251 Phosphorous determination

252 Phosphorous determination results were separated in this study into total (P_{tot}), inorganic (P_{in})
253 and organic (P_{org}) categories. Phosphorus values averaged 788.17 $\mu\text{g/ml}$ P_{tot} , 569.29 $\mu\text{g/mL}$ P_{in} ,
254 and 218.88 $\mu\text{g/ml}$ P_{org} (Tables 4 and 5). The results showed no general trends down-profile
255 (Table 6), but there was a large range in P values represented from weakly to moderately
256 enhanced ($P_{tot} = 224.91\text{-}1300.76$ $\mu\text{g/ml}$ and $P_{in} = 7.85\text{-}892.61$ $\mu\text{g/ml}$). This is in common with the
257 results from previous research done in the Andes, with the exception of the higher values
258 associated with the use of P-rich fertilizer (Eash, 1989; Goodman-Elgar, 2002; Sandor and Eash,
259 1991, 1995). Downslope, profile averages of P_{org} decreased between Rows 6 and 5, from 157.16
260 to 89.77 $\mu\text{g/ml}$, then increased to Row 1 at 402.25 $\mu\text{g/ml}$.

261 Micromorphological analysis

262 Micromorphological analysis was undertaken on samples from the main indicative stratigraphic
263 horizons represented in Test Pits 1, 2, 4 and 5 (see Table 7; Fig. 5). The make-up of the terrace
264 deposits in Test Pit 1 exhibited a poorly sorted, apedal, sandy/silt loam fabric with grains found
265 in all orientations (Fig. 6a). This soil had once contained a greater organic component as
266 indicated by the vughy nature of the soil fabric (Stolt and Lindbo, 2010), but was neither
267 particularly humic nor affected by the secondary formation of amorphous sesquioxides. Similar
268 material continues to be evident down-profile until the weathered tonalite bedrock material is
269 encountered (Fig. 6b). There is no indication of a buried soil present even though this had been
270 hinted at in the field.

271 The terrace make-up in Test Pit 2 was very similar to that observed in Test Pit 1. At the base of
272 the test pit there was a similar sandy/silt loam without much humic or amorphous iron staining,
273 but it did exhibit hints of a small blocky ped structure and occasional aggregates, and a few
274 coatings of pure to dusty (silty) clay in the groundmass (Fig. 6c). This is suggestive of a possible
275 weathered B or Bw (cambic) horizon remnant of a buried soil (Kuhn et al., 2010).

276 Test Pit 3 was not sampled as there was only *c.* 15 cm of present day topsoil over the weathered
277 bedrock. Test Pit 4 was also shallow with only 20 cm of modern topsoil over a possible buried
278 soil that was similar to that in the base of Test Pit 2 (Fig. 6d). The basal horizon of Test Pit 5
279 exhibited a similar fabric to the other possible old land surface in Test Pit 2, a sandy/silt loam,
280 but in this case it had common micro-charcoal and occasional void in-fills of phytolith-rich ash
281 (Fig. 6e). The latter are suggestive of deliberate additions of organic midden-derived material to
282 the soil as fertiliser that have worked their way down-profile in the pore-soil water system and by
283 soil faunal mixing (Stolt and Lindbo, 2010), but do not appear to be a common feature of these
284 terrace soils as observed in the other test pits.

285 The make-up material of the stone terraces of system B on the downhill slopes to the west of the
286 Sangayaico site complex was much thinner than had been expected, ranging in thickness from *c.*
287 56-67 cm. In Test Pit 5 there was a hint of an old land surface present in the basal *c.* 16cm of the
288 profile, but not in Test Pits 21 and 22. The terrace deposits are consistently composed of a poorly
289 sorted mixture of very fine to fine sand-sized quartz and tonalite fragments with a humic silt fine
290 fraction inbetween (Fig. 6a). The thin surviving buried soil/old land surface is composed of a
291 similar fabric but was less porous, somewhat better sorted, with a weakly developed blocky ped
292 structure, occasional pure to dusty clay in the groundmass, and a greater included humic
293 component which also comprised plant derived ash.

294 **Discussion**

295 Soil micromorphological analysis suggests that reasonably well defined old land surfaces/buried
296 soils were only present in Test Pit 2 (and possibly Test Pits 7, 10-17, 19 and 20) in Terrace
297 System A and Test Pit 5 in Terrace System B. Where buried soils were evident, they were thin
298 and patchy and poorly developed, with only minor illuviation of silts and clays evident in the
299 sandy/silt loam.

300 The majority of the *c.* 25-75 cm terrace build-up was composed of a similar sandy/silt loam soil
301 material, but mainly without illuviation features, intermixed with common to abundant tonalite
302 rock fragments of varying sizes. In many respects the terrace build-up material resembles a
303 depleted eluvial Eb horizon, with increased coarser, sand-sized and stone-sized components. The
304 lack of variability in clay content down-profile is likely due to the eluviation of clays from the A
305 horizon (Sandor and Eash, 1995). This points to a combination of lateral and down-profile soil
306 flushing caused by introduced water from irrigation, as well as physical mixing processes
307 associated with past arable use of the terraces, the incorporation of organic matter and strong soil
308 faunal activity, and exposure and weathering of the tonalite bedrock in the upper part of each
309 terrace. It should also be noted that in the Andean sierra, soils are exposed to diurnal freeze-
310 thaw variations that when combined with intense solar radiation and dramatic differences in
311 seasonal and annual variability in rainfall, accelerate the soil mixing processes and the
312 weathering of the parent material and downslope erosion processes throughout the soil system
313 (Contreras, 2010; Goodman-Elgar, 2008; Van Vliet-Lanoe, 2010).

314 The basal terrace soils are essentially stabilised versions of the terrace make-up material above.
315 Any real presence and depth of older (earlier Holocene) soils are ostensibly missing beneath the
316 terrace systems investigated, but geoarchaeological investigations by C French as part of the

317 same overall project have discerned argillic fine sandy clay loam soils present in the Olaya
318 valley about 200m and 2km upstream of Sangayaico. Thus it is possible that these argillic soils
319 (or luvisols) were once more widespread in the catchment, but have generally changed beyond
320 recognition quite rapidly, first to colluvial sandy loams and then to terrace accumulations of
321 rubbly sandy/silt loam over shallow, weakly developed, often truncated, A-B/C or A-B-B/C/C
322 cambisol or leptosol-type soils. Down-slope erosion and associated soil truncation prior to the
323 establishment of the terraces would have been a real consideration in causing this soil change,
324 but are almost impossible to quantify, and it is impossible to rule out previous agricultural
325 activities on the slopes also contributing to this apparent major change in soil type and its
326 survival.

327 The chronology of these changes is much harder to ascribe with any real accuracy. Certainly
328 other examples of Andean terrace systems are fully developed by about 1300 years ago (Beach et
329 al., 2015), and it is reasonable to assume that the terrace systems on the slopes adjacent to
330 Sangayaico are at least associated with the settlements that are dated there to cal AD 1122-1527.
331 Either way, there is a strong probability that the soil-scape on the hillsides has been highly
332 modified by the past establishment of the terrace system(s), perhaps over no more than the past
333 800-900 years or so. The whole soil complex is relatively young and under-developed.
334 Nonetheless, these terrace soils appear to have been well managed, essentially through the
335 repeated addition of organic matter.

336 Arable cropping would have continued to deplete the nutrient and organic matter levels of these
337 terrace soils. This would have occurred despite the regular introduction of water carried down
338 valley along-slope by the main stone irrigation channels, fed by spring/river water from the
339 pampa zone and sluices letting the water downslope into each set of terraces (Denevan, 2001),

340 the continuing deliberate addition of organic matter from the turning in of harvested crops,
341 pastoral herds being kept on these fields in-between cropping seasons and any deliberate
342 additions of household midden debris. To the detriment of the wider soil system on the valley
343 sides, irrigation and rainfall combined would have encouraged the flushing of fines and nutrients
344 from these soils down-profile and down-slope as lateral flushes, possibly counter-acting the
345 moisture retention aspect of the thickened terrace soils themselves. Consequently long fallow
346 recovery periods of several years would have been required to maintain a reasonable fertility in
347 these soils as well as regular burning of the fields after each growing season (as practised today).
348 Even then, recovery of fertility would have been slow and any real soil development unlikely, a
349 feature which is recognised today despite much of this highland area being abandoned and
350 largely unused.

351 In general the physical characteristics revealed in this study of the Sangayaico terrace system do
352 not show the same clear evidence of soil thickening, modification and long-term fertilisation that
353 was observed by similar studies such as in the Colca Valley of Peru (Eash and Sandor, 1995),
354 nor the distinctive increase in clay illuviation noted in the soils of the Tococtocasa terrace system
355 in the Chicha-Soras valley (Kemp et al., 2006; Branch et al., 2007). In the Colca study it was
356 observed that A horizons were commonly thicker by *c.* 30-130 cm, exhibited a lower bulk
357 density (implying a much greater organic content), and the upper horizons were enriched with
358 organic matter. Other studies observed lower pH values, more organic carbon and nitrogen, the
359 addition and inclusion of midden-derived material, and the associated deep translocation and
360 enhancement with phosphorus of the buried B horizons at the base of the terrace profiles
361 (Goodman-Elgar, 2009; Sandor et al., 2007; Wells, 2006). In the Chicha-Soras terrace soil study
362 (Kemp et al., 2006; Branch et al., 2007), the buried upper terrace and surface terrace soils both

363 exhibited an abundance of illuvial clay coatings which were attributed to the weathering,
364 disturbance and down-profile migration of neo-formed clay from the volcanic clasts on site,
365 aggravated by the oscillating arid/humid climate and the repeated input of irrigation water. In
366 contrast, the Sangayaico terraces rarely exhibit over-thickened A horizons, even though there is
367 regularly *c.* 25-75 cm of cumulative terrace soil aggradation. The buried B horizons are either
368 thin or not present, and exhibit little sign of enrichment with illuvial clay (or argillic clay) down-
369 profile. Organic matter, nitrogen and phosphorus values are weakly variable and relatively only
370 weakly enhanced, and we know that local farmers did not use artificial fertilisers over the past 50
371 years or so (from anecdotal accounts of local farmers). This may indicate the lack of recent
372 irrigation as well as a longer-term gradual process of neglect and lack of arable use and
373 fertilisation, and general degradation through hillwash and lateral flush effects through these
374 terrace slopes. Again this is in sharp contrast to the Colca valley study where the farmers appear
375 to have known the exact state and characteristics of their land and how to improve, conserve and
376 husband it successfully (Sandor and Furbee 1996).

377 In this Viejo Sangayaico study, no conclusive indicators were observed to more precisely
378 indicate which agricultural strategy may have been employed. The paucity of ceramic and faunal
379 remains and charcoal in and on the terraces themselves, often associated to the removal of
380 household waste (Goodman-Elgar, 2002), and the lack of spikes in soil organic matter and
381 phosphorous (P_{tot} , P_{in} , or P_{org}) averaged across rows, would indicate that midden material was not
382 generally added to the terrace surfaces. Without a natural soil profile to compare to as a control,
383 it is hard to provide quantitative support for an argument attesting to the extent of manuring. But
384 given the similarities of the results presented here to that of Homburg et al. (2005), and the lack
385 of extraordinary peaks in phosphorus as discussed in Sandor and Eash (1991), it is likely that the

386 terrace systems associated with Sangayaico were never extensively fertilised through manuring.
387 Based on ethnographic evidence and field observation, it is probable that manuring was mainly
388 done by grazing animals following the harvest or during a fallow period (De la Vega, 1960;
389 Guillet, 1981, 1987; Zimmerer, 1998). Finally there is no conclusive evidence for seasonal field
390 burning in the Sangayaico terrace systems. Indeed, there was very little charcoal found in the
391 terrace profiles, except for the uppermost levels of the A_p horizon in Terrace System B, which
392 probably indicates a recent burning.

393 As to whether past agricultural strategies have affected long-term soil fertility, the terraces
394 associated with Sangayaico showed little evidence for degradation. But, substantial fallow time
395 and the grazing of animals on the crop stubble would have helped ameliorate this system.
396 Anecdotal conversation with the local farmers working around Sangayaico suggests that the
397 fields remain productive without the use of artificial fertilisers due mainly to the use of a five-
398 year fallow period following a two-three-year growing period. Another important factor in the
399 preservation of the Sangayaico agricultural landscape is the continued use of the *chaquitacla*, or
400 Andean foot plough, to turn the soil in the fields. Indeed, using the *chaquitacla* greatly reduces
401 the breakdown of beneficial soil aggregates as opposed to mechanised ploughing (Goodman-
402 Elgar, 2008).

403 It should also be noted that Terrace System A was selected because it appeared to be one of the
404 best preserved terrace systems associated to the Sangayaico site, whilst other terrace systems in
405 the vicinity have fallen into disrepair. The resulting differences between denuded areas and those
406 with better-preserved terracing provides an example of the cost of abandoning or neglecting
407 intensive agricultural systems. Such abandonment and neglect was widespread throughout the
408 New World with the arrival of the old-world diseases and the relocation of indigenous

409 populations to live in *reducciones*, and then again in the last thirty-five to fifty years with
410 massive sierra-to-coast population shifts (Denevan, 2001; Donkin, 1979; Fisher et al., 2003;
411 Gade and Escobar, 1982; Wernke, 2010).

412 **Interpretative discussion and conclusions**

413 This study has provided the first characterisations of the agrarian soil system in the upper Ica
414 valley surrounding Viejo Sangayaico. The results suggest the relative stability of the terraced
415 systems themselves over the past millennium, but major transformations of the underlying old
416 land surface had already occurred prior to the establishment of the terraced field systems still
417 visible today. Although it requires further investigation and reliable dating, it is very possible
418 that it was erosion, mixing and depletion associated with earlier, pre-Late Intermediate period (or
419 pre- *c.* AD 1100) land-use (for both arable agriculture and pastoralism) that had caused such
420 major soil change on the hill-slopes around Sangayaico. In contrast over the last 900 years or so,
421 local agriculturalists were able to sustainably farm the landscape through the construction of
422 irrigated terraces and the use of crop cycles dependent on long-fallowing. This appears to be
423 largely without the extensive use of fertiliser, in contrast to observations in the Colca Valley
424 (Eash and Sandor, 1995; Sandor and Eash, 1991, 1995).

425 It is suggested that the populations associated with sites throughout the upper and middle Ica
426 River drainage relied on terraces on the mountain slopes as a means of insuring soil stability and
427 conservation. Terraces controlled both landscape erosion and degradation, thus increasing the
428 production area and creating the advantages of a local micro-climate which provided a growing
429 area more amenable to crops normally grown at lower elevation. It is possible that with the
430 arrival of the Spanish and subsequently into the Republican period, the relocation of villagers
431 such as the Sangayaiqueños to mines in the Colonial Period (Bueno Mendoza, 2003; Maldonado

432 Pimentel and Estacio Tamayo, 2012) and subsequent emigration to urban cities (Zimmerer,
433 1991), the terraces began to fall into disrepair, causing an increase in slope instability and
434 concomitant soil erosion.

435 Given the relatively low and variable organic matter content and minimal plant macro-nutrients
436 such as phosphorus in the terrace soil systems around Sangayaico, experienced learned
437 knowledge would have been needed to select the appropriate best crop for each area. The amount
438 of forethought in crop selection would have called for a great deal of knowhow about the
439 behaviour of local crops and soils, as has been observed by Sandor and Furbee (1996) for the
440 Colca Canyon. By acknowledging that the conditions for plant growth do not necessarily occur
441 within strictly delimited ecozones or crop range limits but in a more complex mosaic across the
442 landscape, in effect the range of crops can be ‘stretched’ (Zimmerer, 1999).

443 Thus strict adherence to models relying on defined ecotones appears to be ill-advised and that
444 using Zimmerer’s (1999) “overlapping patchwork” framework approach appears to offer a more
445 holistic and comprehensive explanation of Andean ecology and agricultural land-use. Although
446 the relative instability of the underlying soil properties in this region may hinder specialised
447 cropping, variation in erosion factors would have favoured mixed-cropping systems and allowed
448 a more diverse range of crop production (Zimmerer, 1999). Future palynological studies of the
449 vegetational sequences in the basin mires in the puna up-valley from Sangayaico may shed better
450 light on this suggestion of mixed cropping in due course. Nonetheless, this variability is well
451 documented in Andean soil systems from previous studies (Eash, 1989; Eash and Sandor, 1995;
452 Goodman-Elgar, 2002; Goodman-Elgar, 2008; Kemp et al., 2006; Sandor and Eash, 1991, 1995)
453 as well as at the landscape scale through ecological and human land-use observations (Contreras,
454 2010; Branch et al., 2007; Zimmerer, 1999). The main reason for this variability is due to the fact

455 that the Andean agro-ecosystem is the result of the coupling of localised climatological,
456 geological and geomorphological processes (Montgomery et al., 2001), shaped, altered and
457 managed by tremendous human endeavour.

458

459 **Acknowledgements**

460

461 The authors would like to thank our hosts, benefactors and collaborators in Peru, particularly
462 Alberto Benavides, Susanna Torres Acres and George Chauca. Funding for the field and
463 laboratory work was provided by the Leverhulme Trust, a private donation, and the Department
464 of Anthropology, Washington State University. Many thanks to Professors Melissa Goodman-
465 Elgar, James Harsh, and Timothy Kohler for advising on the soil fertility and archaeological
466 research aspects completed at Washington State University. We would also like to thank Tonko
467 Rajkovaca of the McBurney Geoarchaeology Laboratory, Department of Archaeology and
468 Anthropology, University of Cambridge, for making the soil thin sections, as well as Jeffery
469 Boyle and Margaret Davies of the Department of Crop and Soils Sciences, Washington State
470 University, for their assistance with the soil fertility analyses.

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690

691 **Figure List**

692

693 1. Location map of Sangayaico in the upper Ica valley of southern Peru (D. Beresford-Jones,
694 based on LANDSAT 7 ETM+ 2000, USGS)

695 2. The location (left; based on Google Earth: Image@ 2014 Digital Globe/@2014 Cnes/Spot
696 Image) and schematic plan of rows/numbers of test pits (right) in Terrace Systems A and B (red;
697 B, blue). Note that test-pit locations are white shapes and the location of Viejo Sangayaico site is
698 circled in yellow. (W. Nanavati)

- 699 3. A view of terrace system B today looking up-slope to the Viejo Sangayaico sites (C. French)
- 700 4. Terrace wall and profile photograph of Test Pit 1 in terrace system A (C. French)
- 701 5. Selected sections of TP1, 13 and 18 in Terrace System A, and TP5, 21 and 22 in Terrace
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- 705 Test Pit 1, sample 3/1 (frame width 4.5mm; cross polarized light)
- 706 b. Photomicrograph of weathered B/C of tonalite fragments, Test Pit 1, sample 4 (frame width
- 707 4.5mm; cross polarized light)
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- 712 e. Photomicrograph of the phytoliths in the ash infill, Test Pit 5, sample 5/2 (frame width
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- 714 **List of Tables**
- 715 1. Summary of methods used to discern agricultural inputs (OM= organic matter; Na-Hex=
- 716 sodium hexametaphosphate)
- 717 2. Field description of the Test Pits at Sangayaico (Note: Profiles 1-4 and 6-20 were taken from
- 718 Terrace System A; Profiles 5, 21 and 22 were taken from Terrace System B)
- 719 3. Analyses used in this study
- 720 4. Summary of the bulk physical results for pH, electrical conductivity, soil moisture content,
- 721 organic/inorganic carbon, total carbon and nitrogen, soil texture and phosphorus

- 722 5. Bulk sample results for pH, C:N ratio, particle size distribution and phosphorus
- 723 6: Trends in soil characteristics down-profile and down-slope for each terrace system
- 724 7. Summary micromorphological descriptions and interpretations