¹ Capillary Trapping of CO₂ in Oil Reservoirs:

² Observations in a Mixed-wet Carbonate Rock

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7 KEYWORDS

8 Residual trapping, capillary trapping, remaining saturations, mixed-wet, supercritical CO₂, CCS,

9 EOR, X-ray, carbon utilization.

10

11 ABSTRACT

Early deployment of carbon dioxide storage is likely to focus on injection into mature oil reservoirs, most of which occur in carbonate rock units. Observations and modeling have shown how capillary trapping leads to the immobilization of CO_2 in saline aquifers, enhancing the security and capacity of storage. There are, however, no observations of trapping in rocks with a mixedwet state characteristic of hydrocarbon bearing carbonate reservoirs. Here we found that residual trapping of supercritical CO_2 in a limestone altered to a mixed-wet state with oil was significantly

less than trapping in the unaltered rock. In unaltered samples, the trapping of CO₂ and N₂ were 18 19 indistinguishable, with a maximum residual saturation of 24%. After altering the wetting state, the 20 trapping of N₂ was reduced, with a maximum residual saturation of 19%. The trapping of CO₂ was 21 reduced even further with a maximum residual saturation of 15%. Best fit Land model constants 22 shifted from C = 1.73 in the water-wet rock, to C = 2.82 for N₂, and C = 4.11 for the CO₂ in the 23 mixed-wet rock. The weakened trapping indicates that plume migration will be farther and the timescales for immobilization will be longer for CO₂ storage projects using oil fields compared 24 25 with saline aquifers.



28 INTRODUCTION

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The capture of carbon dioxide emitted from industrial processes and subsequent storage in subsurface geologic units has been identified as a key technology in the global reduction of anthropogenic CO_2 emissions to the atmosphere¹. Capillary trapping, also called residual trapping, is one of the most significant physical mechanisms ensuring the permanence of CO_2 storage in the subsurface^{1,2}. Capillary trapping limits the extent of CO₂ migration^{3,4}, underpins
estimates of regional storage capacity⁵⁻⁸ and controls the movement of the mobile sections of a
subsurface plume through its relationship with fluid flow property hysteresis^{4,9,10}. Thus a
significant body of recent work has focused on observations characterizing capillary trapping²,
including in carbonate rocks¹¹.

Saline aquifer storage presents the largest potential for global CO_2 storage capacity¹² and the 38 39 focus of the previous observations has been on the characterization of rocks unaltered by 40 hydrocarbons. On the other hand, oil fields have a number of characteristics that could lead to 41 their having an outsize importance during the early deployment of carbon storage. Commercial 42 oil fields are well characterized, have a demonstrated fluid trap, and sometimes have infrastructure that can be repurposed for use with CO₂ injection^{13,14}. Enhanced oil recovery with 43 44 CO₂ injection may also provide for a significant revenue stream. These benefits are reflected in 45 the dominance of enhanced oil recovery in the current portfolio of existing industrial scale sequestration projects¹⁵. 46

47 The capillary properties of rocks typical of saline aquifer systems - unaltered by hydrocarbons 48 - are considered water-wet with respect to CO₂-brine systems across wide range of reservoir 49 conditions¹⁶, although some important questions remain including the impact of chemical process kinetics on the wetting state^{17,18}. In contrast, hydrocarbon reservoirs are characterized by 50 51 a mixed-wet state, wherein chemical deposits from the hydrocarbon phase onto the surface of the 52 pore mineral surfaces over geologic time have resulted in connected oil-wetting conduits to 53 flow¹⁹. This has a significant impact on the flow properties of the system, including the capillary 54 pressure, relative permeability, and residual trapping characteristic functions²⁰⁻²².

55	Of key importance for the security of CO ₂ storage is that the capillary trapping of nonpolar
56	fluids, like CO ₂ , have been observed to be significantly less in mixed wet rocks than trapping in
57	water wet rocks typical of saline aquifers ²³⁻²⁵ . This issue has been largely ignored in the flow
58	modelling literature for CO ₂ storage and there are no observations characterizing the extent of
59	capillary trapping that will take place with CO ₂ in mixed-wet rocks. In this work we provide the
60	first observations that a key trapping mechanism underpinning the security of CO ₂ storage is
61	significantly reduced in the very storage locations that are among the most economically
62	appealing for the first generation of project development.
63	The key aim of this study was to evaluate the impact of the wetting state typical of carbonate
64	oil reservoirs on trapping for CO ₂ storage. We have done this by characterizing the trapping of
65	supercritical CO ₂ in a mixed-wet carbonate rock, comparing this with the degree of trapping in
66	the same rock prior to alteration to the mixed wet state. We focused on a carbonate rock because
67	of the prevalence of carbonate lithology in oil production ^{26,27} . To isolate the effects of wetting
68	alteration from fluid-rock chemical reaction and mass transfer, the CO ₂ and brine were
69	equilibrated together with crushed samples of the rock prior to observations of trapping.
70	Experiments were also performed with N_2 –water systems in unaltered and altered samples as a
71	benchmark for comparison.

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73 MATERIALS

Rock Samples, Fluids, and Test Conditions. A total of 14 core flood experiments were
 carried out on two Estaillades limestone core samples before and after alteration with a

hydrocarbon mixture. The sequence of experiments, fluid pairs, and drainage flow rates areprovided in Table 1.

Estaillades is a quarry carbonate limestone from the Estaillade Formation located in southeast France²⁸. The sample consists of 97.9% calcite and 2.1% quartz, measured by Weatherford Laboratories, East Grinsted, UK. The faces of the core were machined flat to ensure contact with the end-caps. Each sample was vacuum dried at 70 °C overnight before each test.

82 A single sample, E1, was used to compare trapping between water-wet and mixed-wet 83 systems. Observations were first made on the sample unaltered by hydrocarbon, and again after 84 altering the wetting state to a mixed-wet state using a procedure described in more detail below. 85 A second sample, E2, was used for a reproduction of observations in the unaltered rock. Both 86 samples were 3.8 cm in diameter and 12 cm in length. The absolute permeability of the E1 and 87 E2 samples to water were 138 and 148 mD, respectively. The average porosities as measured 88 with X-ray computed tomography (CT) were 28 and 29% for the E1 and E2 samples, 89 respectively.

90 Carbon dioxide and nitrogen were used as the non-polar fluid phases in the core-flooding 91 experiments, both with 99.9% purity (BOC Industrial Gases, UK). Nitrogen was used as the 92 benchmark nonwetting phase primarily because the thermophysical properties, including the 93 wetting state in unaltered carbonate rocks, was well constrained. There were other considerations 94 - the viscosity of N₂ is more similar to that of CO₂ than a typical benchmark hydrocarbon liquid 95 like decane. The use of N₂ allowed us to perform experiments in sequences alternating fluid pairs 96 without the need for a solvent based cleaning of the rock core, a process which may have altered 97 the wetting state and put further stress on the rock. The aqueous phase fluids used were

deionized water or brine consisting of deionized water and NaCl with total salt molality of 1 mol
kg⁻¹. Arabian Medium crude oil (API=30.77°) and heptane with 99% purity (Sigma-Aldrich)
were used for the wettability alteration and Amott tests²⁹. Toluene with 99.8% purity (SigmaAldrich) was used as a solvent in the filtration experiments made prior wetting alteration to
measure asphaltene precipitation of the crude oil.

103 All core flooding experiments were performed at 10 MPa pore pressure. The CO₂-brine

104 experiments were performed at a temperature of 50 °C while N₂-water experiments were at 25

105 °C. The thermophysical properties of the fluids are provided as follows: The interfacial tension

106 value of CO_2 and brine at the experimental condition was estimated to be 37 mN/m³⁰ while N_2

107 and water interfacial tension value was 67 mN/ m^{31} . The CO₂/brine and N₂/water viscosities were

estimated to be 0.03/0.61 and 0.02/0.89 cP, respectively^{32,33}. The CO₂/brine and N₂/water

109 densities were estimated to be 0.39/1.02 and 0.11/1.01 g/cm³, respectively³³.

Table 1. Initial-residual core flooding experiments and repeats in the chronological sequence
that they were performed.

Experiment number ^a	Sample name / wetting state	Fluid pairs	Drainage flow rate ^b [ml min ⁻¹]	Pore volumes injected during imbibition
1	E1 / water-wet	N_2 / water	0.5	0.25
2	E1 / water-wet	N_2 / water	20	0.17
3	E1 / water-wet	N_2 / water	_c	0.63
4a	E1 / mixed-wet	N_2 / water	20	0.2
5a	E1 / mixed-wet	N_2 / water	0.5	0.36
6a	E1 / mixed-wet	CO_2 / brine	20	0.2
7	E1 / mixed-wet	CO_2 / brine	0.5	0.63
4b	E1 / mixed-wet	N_2 / water	20	1.29
5b	E1 / mixed-wet	N_2 / water	0.5	1.71
8	E2 / water-wet	N_2 / water	20	0.45
9a	E2 / water-wet	CO2 / brine	20	0.44

9b	E2 / water-wet	CO_2 / brine	20	0.98
10	E2 / water-wet	CO ₂ / brine	0.5	0.23
6b	E1 / mixed-wet	CO_2 / brine	20	0.23

^a Experiments with a common number indicate repeat tests on the same sample, at the same wetting state, and fluid flow condition, i.e., test 4b is a repeat of test 4a, but after the rock core has been exposed to CO_2 during tests 6a and 7.

- ^b Imbibition was at 0.5 ml min⁻¹ for all experiments
- 116 ^c Experiment 3 was initially fully saturated with N_2 instead of water. Water was subsequently injected at 0.5 ml 117 min⁻¹ to measure the initial-residual relationship at 100% initial gas saturation

118 **Core-flooding Experimental Setup.** Residual trapping measurements were conducted using a 119 two fluid phase core flooding system described in detail in previous papers by the authors¹⁶. The 120 experimental setup was designed to maintain temperature and pressure conditions of up to 120 °C 121 and 30 MPa. The system was modified for experiments with carbonates samples and included a 122 stirred reactor to equilibrate fluids with ground sample of the solid (Parr Instruments Co., IL, 123 USA). A schematic is provided in the supporting information.

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125 METHOD

126 **Core-flooding Test and Sequence.** A core flooding technique³⁴ was applied in this work for 127 the construction of initial-residual characteristic trapping curves. The technique combines the 128 deliberate creation of capillary pressure gradients in the core with in-situ saturation monitoring to 129 rapidly construct the initial-residual curve across a range of saturations. A detailed description of 130 the startup and operation of the experiment is provided in the supporting information and 131 summarized here.

132 To create the initial saturation, a primary drainage step was performed, injecting CO_2 or N_2 at a 133 constant rate of either 0.5 or 20 ml min⁻¹ into an initially water or brine-saturated core sample, 134 Table 1. The choice of flow rate was made at different stages in the experiment to cover a particular 135 range of the initial-residual curves – larger flow velocities during drainage result in larger initial 136 saturations. X-ray scans were then performed to evaluate the initial saturation along the length of 137 the rock core. Imbibition was then performed by injecting brine or water, pre-equilibrated with CO₂ and the limestone sample, at 0.5 ml min⁻¹. Saturations were calculated at each stage with X-138 ray CT using the standard approach³⁵ detailed in the supporting information. Several scans were 139 140 taken to confirm measurements of imbibition and to reduce the uncertainty associated with the 141 computed saturations.

142 To minimize the impact of calcite in the rock dissolving into the acidified brine, the brine was 143 pre-equilibrated with both CO₂ and the rock material at the experimental conditions prior to the 144 residual trapping test. The fluids were mixed vigorously together with crushed grains of the same 145 rock in a Parr stirred reactor at experimental conditions for 24 hrs. Then, this CO₂ and brine were 146 co-circulated at experimental conditions in a closed loop bypassing the core, but including 147 circulation through the reactor. The dissolution of CO_2 into water was monitored by observing the 148 volume balance of the fluid phases in the closed loop, maintained at constant pressure. X-ray 149 imaging was used to confirm that the porosity was not changing during the tests. A comparison of 150 the repeatability of the results of N₂ tests before and after CO₂ flooding tests also supported the 151 view that the rock core, and its wetting state, were not significantly altered by the exposure to CO₂. 152 The absolute permeability to water was measured before each experiment and varied less than 4% 153 throughout the series of tests suggesting minimal changes in the pore structure of the core sample 154 throughout the entire course of experiments. The slice average porosity along the entire length of 155 core was also measured at a resolution of 1 mm after each test with X-ray scanning providing156 another confirmation that rock properties were not significantly altered.

157 The sequence of tests were designed to test the impact of altering the wetting state and confirm 158 that variations in observations were due to the change in wetting state. See Table 1. Tests 1-3 were 159 used to construct the initial-residual curve for sample E1 using N₂ with the rock in the natural state, 160 unaltered by hydrocarbon. The wetting state was altered as described below and subsequent tests 161 characterizing the initial-residual curve were performed first using the N_2 -water system (tests 4a, 162 5a) and then the CO_2 -brine system (tests 6a, 7). Repeat tests were then performed with N_2 (tests 163 4b, 5b) to assess the impact, if any, of CO_2 exposure to the altered wetting state. A final repeat test 164 was performed with CO_2 (6b) to test repeatability of the residual trapping at high initial saturation 165 with CO₂.

At the outset of the study, we hypothesized that the N₂-brine and CO₂-brine initial-residual curves would match based on previous experience³⁴. As will be discussed in the results, this turned out not to be the case for the initial residual curves obtained after altering the rock with hydrocarbon. To ensure that this effect was a result of the wetting state change, and not the propagation of a feature of CO₂ and N₂ trapping in unaltered carbonates, another set of tests was performed on a second sample of unaltered Estaillades rock comparing N₂ and CO₂ trapping, tests 8-10.

173 Injected pore volumes were kept to a minimum due to the corrosive nature of the fluids and the 174 relative fragility of the rock under the test conditions. This is a particular difficulty of scientific 175 studies using reservoir condition core floods – a single rock sample must endure a sequence of 176 tests which place significant stress on the rock. In water wet systems, a single pore volume is

sufficient to achieve the residual saturation³⁴, whereas hundreds or even thousands of injected pore 177 178 volumes are required to achieve the true residual in mixed-wet systems²⁴. In both water-wet and 179 mixed-wet systems there is an initial rapid desaturation of the non-wetting phase with less than 180 one pore volume of fluid injected. In the mixed-wet case, however, this is followed by a very slow 181 rate of desaturation during subsequent brine injection. In the tests performed in this work, the 182 initial rapid desaturation was monitored using X-ray scanning. The point of comparison of the 183 initial-residual curves was made using saturation measurements obtained following this initial 184 phase of rapid desaturation, generally between 0.25 and 1.5 pore volumes of brine injected, Table 185 1. The saturation of CO₂ in a mixed-wet system after such volumes of imbibition will be similar 186 to saturations in a field setting far from the well where the CO_2 has been injected. However, the 187 ultimate residual in the mixed-wet systems, should they ever be obtained, will be lower than the 188 reported values obtained in this work.

189 Altering the Wetting State of the Rock. After completing observations with the unaltered E1 190 sample, an aging technique using crude oil was applied to change the wetting state of the core sample. An approach similar to Salathiel 1973²⁴ was performed by preparing a mixture of the 191 192 evacuated crude oil with an organic precipitant, heptane. The addition of heptane induces the 193 precipitation of asphaltene, leading to the alteration of the wetting state of mineral surfaces. Before 194 aging the rock, observations were made with crude oil-heptane mixtures to identify ratios of fluid 195 resulting in sufficient asphaltene precipitates to induce the wetting state change, but not so much 196 that permeability of the rock core is significantly reduced. A detailed description is provided in the 197 supporting information.

198 The core sample was fully saturated with evacuated brine and heated to 70 °C. A freshly made 199 oil mixture of 28% evacuated Arabian Medium crude oil and 72% heptane was heated to 70 °C

200 and injected into the core sample for a total of 20 pore volumes with the direction of flow reversed 201 midway to dynamically age the sample and speed the wetting alteration process³⁶. The flow rates ranged from 0.1 ml min⁻¹ and up to 5 ml min⁻¹. The sample was then left for further static aging 202 203 (no flow) at 70 °C for 40 days. Then, the sample was cooled down to room temperature for at least 204 24hrs. The oil mixture was displaced with 5 pore volumes of heptane injected at 10 ml min⁻¹ at 205 room temperature. The core sample was then vacuum dried at 100 °C for three days to remove 206 heptane and brine from the pore space by evaporating the fluids (heptane and water). Then the core 207 sample was cooled to room temperature. Degassed deionized water was flushed into the sample 208 for several pore volumes to remove salt that might have precipitated in the sample during drying. 209 The sample was vacuum dried again at 70 °C for 12 hours to remove water. Finally, to ensure mobile oil was completely displaced, 5 pore volumes of CO₂ was injected at 20 MPa³⁷. Next we 210 211 performed the initial-residual core flood tests.

212 After completing the initial-residual core flooding experiments, the wetting properties of the 213 mixed-wet (E1) and water-wet (E2) samples were measured by carrying out a standard Amott test²⁹ using Heptane and brine. Details of the test are provided in the supporting information. The 214 215 Amott water ratio wettability index, δw , was measured to be $\delta w = 0.48$ and $\delta w = 0.89$ for E1 and 216 E2, respectively, after nine pore volumes of brine injected. The high δw value of E2 is indicative 217 of a strongly water-wet system where spontaneous imbibition leads to most of the oil in the sample 218 displaced by water. On the other hand, spontaneous and forced imbibition produced comparable 219 volumes of oil in the E1 sample resulting in a wettability index value similar to those measured in giant carbonate oil fields in the Middle East characterized as mixed-wet systems³⁸. Additionally, 220 221 the water-wet sample (E2) produced little oil after water breakthrough and quickly reached residual 222 oil saturation. On the other hand, in the mixed-wet sample (E1) oil production continued for the entire nine pore volumes of water injected, and likely would continue with continued brine injection. The displacement followed the pattern of a rapid desaturation with less than a single pore volume of brine injected, followed by a slow rate of desaturation with continued brine injection.

Trapping Model. A large number of models have been developed for describing the initialresidual relationship³⁹. The Land model was used to fit the trapping data. It is one of the earliest and most widely used trapping models⁴⁰. The residually trapped non-wetting phase saturation after imbibition, $S_{CO_{2,r}}$ is an increasing function of the maximum non-wetting phase saturation achieved during drainage, $S_{CO_{2,i}}$,

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$$S_{CO_2,r} = \frac{S_{CO_2,i}}{1 + CS_{CO_2,i}},$$
(1)

where *C* is a constant known as the Land coefficient and is equal to or greater than zero. Larger values of the constant indicate less trapping. Some models, e.g., the model of Spiteri et al. $(2008)^{41}$ have been proposed specifically for mixed-wet systems. In this case we found the Land model to fit our observations well. The data can in principle be fit using any of the alternatives.

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238 RESULTS

Figure 1 shows examples of the slice average saturation along the rock core with corresponding two dimensional saturations maps of a central slice of the core for three of tests with high drainage flow rates. These graphs were generated for each experiment and used to quality check the results. The saturation maps show that CO₂ and N₂ were uniformly distributed in the core sample in central

243 parts of the core where the residual trapping was characterized, and not affected by mass transfer 244 or gravity segregation. The slice averaged saturations of the water-wet systems showed a declining 245 trend towards the outlet due to the capillary end effect. On the other hand, the saturation profiles 246 in the same rock sample after wetting state alteration exhibited an end-effect behavior, either with 247 no decline or increasing non-wetting phase saturation, characteristic of the mixed-wet state. The 248 drainage profiles were relatively uniform and saturation increased towards the outlet during 249 imbibition. Data from the central ~60 mm of the rock core were used to construct the initial-250 residual curves to avoid the interference of the boundary affects.



Figure 1. 2D saturation maps of the central slice of the core after drainage (top) and imbibition (middle) for three experiments on the same sample. The core average saturation of the final section of the core sample is shown in the lower graphs. The end-effect changes after the wettability alteration.

In previous work it was confirmed that residual trapping in the water wet case was obtained rapidly, with less than a pore volume of brine injected, and remained stable for more than 40 pore volumes of subsequent brine injection³⁴. In this work the initial rapid desaturation was again observed in tests with both the water wet and hydrocarbon altered rock. The subsequent stability of the residual saturation in the water-wet rock, and the slow and persistent desaturation of nonwetting phase in the altered rock, was then confirmed with the forced injection part of the Amott test shown in Figure 2.



Figure 2. Results from the forced imbibition part of the Amott test showing oil saturation as a function of the pore volumes (PV) of brine injected, applied after 30 days of spontaneous imbibition. The water-wet sample (E1) produced little oil after water breakthrough while the mixed-wet sample (E2) continued oil production with a reducing remaining oil saturation trend as a function of PV injected.

Figure 3 shows initial-residual characteristic curves measured in this study with best fit curves using the Land model, Equation 1. Table 2 shows the best fit parameters for the Land hysteresis model. The trapping of CO_2 and N_2 in the unaltered samples, Experiments 1-3, 9-10, were indistinguishable, with best fit Land constants equivalent within the uncertainty bounds, Table 2.

This was consistent with earlier work³⁴ characterizing CO₂ trapping in unaltered sandstone rocks and suggests that the wetting state is similar for the two systems¹⁶, but contrasts with observations made using a different technique in which more trapping was observed for supercritical CO₂ relative to gaseous CO₂ in an unaltered limestone¹¹. The best fit Land constant for all of the data (both CO₂ and N₂) in the water wet tests was C = 1.73. At the highest initial saturations observed, 42% of the pore space, the residual saturation obtained was 24%.



Figure 3. Initial and residual CO_2 and N_2 saturations measured in this study. White symbols are observations from the unaltered rock. Grey symbols are observations of N_2 -water from the mixedwet rock. Black symbols are observations of CO_2 -brine from the mixed-wet rock. The solid lines show the best fit Land trapping model with their respective values for the parameterization

284 constant, *C*. The order of the experiments in the legend shows the sequence in which the 285 observations were made.

286 After alteration of the wetting state of the rock sample with the crude oil-heptane mixture, less 287 trapping was observed with both N_2 and CO_2 , Experiments 4-7. The value of the best fit Land 288 constant increased from 1.68 to 2.83 for N_2 after 0.2 - 1.7 pore volumes of brine injected into the 289 core. The initial-residual relationship for N_2 was obtained in repeated tests before (Experiments 4a 290 and 5a) and after (Experiments 4b and 5b) tests with CO₂. The data were indistinguishable, with 291 Land constants $C = 2.86 \pm .06$ and $2.80 \pm .06$ from the datasets obtained before and after exposure 292 of the altered rock to CO₂, respectively. This suggests that exposure to CO₂ did not alter the wetting state of the rock sample. 293

The trapping of CO₂ was reduced even further compared to N₂. The best fit Land constant for CO₂ increased to a value of 4.11 after 0.2 - 0.6 pore volumes of brine injected, with slow desaturation continuing up to 1.6 pore volumes of injected brine (see the supporting information). At the measured initial saturation of 42% the remaining CO₂ was reduced to 15% compared with 19% for the N₂ system. This observation was repeated with experiment 6b after exposure to N₂, showing both the precision in the measurement and that the rock core was not substantially altered through multiple measurements obtained in the same sample.

Table 2. Best fit Land residual trapping hysteresis model coefficient

Experiment number	Wetting state	Trapped gas	С*
1, 2, 8	water-wet	N_2	1.68 ± 0.05
9a, 9b, 10	water-wet	CO_2	1.74 ± 0.05
1, 2, 8, 9a, 9b, 10	water-wet	CO_2, N_2	1.73 ± 0.03
4a, 4b, 5a, 5b	mixed-wet	N_2	2.83 ± 0.04

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6a, 6b, 7 mixed-wet CO_2 4.11 ±0.08 * *C* values ± error calculated as two times the standard deviation of best fit *C*

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304 DISCUSSION

305 The results for the unaltered rocks show the same degree of trapping for CO₂ and N₂. The Land 306 constant is larger, indicating less trapping, than values obtained characterizing CO₂ trapping in 307 Berea sandstone². The source of the difference may be due to pore structure and/or wetting state 308 differences², although the similarity of the N₂ and CO₂ observations in this study suggests pore 309 structure differences to be the likely explanation. The equivalence of the CO₂ and N₂ is consistent 310 with past observations with Berea sandstone³⁴, but contrasts with observations made in a different 311 carbonate rock in which more trapping was observed for supercritical CO₂ as compared with gaseous CO_2 in a different carbonate rock¹¹. The contrasting results could be due to differences in 312 313 rock structure leading to different sensitivities to contact angle changes, or experimental 314 procedure, and is discussed further below.

In the mixed-wet sample the trapping of N_2 and CO_2 systems were significantly less than the trapping in the unaltered rocks. As discussed in the introduction, this result was consistent with past observations evaluating the impact of the mixed-wet state on the capillary trapping of hydrocarbon liquids^{23-25,42}. This is generally thought to result from oil-wet surfaces providing connected pathways for thin film drainage of non-polar fluids^{24,40}.

320 The trapping was also specific to the fluid pair, with less CO_2 remaining in the pore space than 321 N₂ upon secondary imbibition. The observations with N₂ and CO_2 were alternated and repeated 322 with the same result, indicating that the difference was not due to chemical reaction between the fluids and the rock, or altered wetting surfaces. This suggests that the difference was due to the impact of a difference in contact angle between the CO_2 -brine and N_2 -water on the surfaces with an altered wetting state, with the CO_2 more wetting of the altered surfaces than N_2 .

This response is predicted by theoretical considerations of contact angles on smooth surfaces⁴³. 326 327 With a contact angle, θ [rad], measured in the aqueous phase, the relationship between the CO₂-328 brine and N2-brine angles are dependent on the ratio of their respective values of interfacial tension, σ , [mN/m], through the relationship $\cos \theta_{CO_2-w} \propto \frac{\sigma_{N_2-w}}{\sigma_{CO_2-w}} \cos \theta_{N_2-w}$. A derivation is included in 329 330 the supporting information. The interfacial tension of N₂ and gaseous CO₂ systems are much larger than that of the supercritical CO₂ system^{30,31}, predicting a larger contact angle for CO₂ (more CO₂-331 332 wetting) on oil-wet surfaces and smaller contact angles for CO₂ (less CO₂-wetting) on water-wet 333 surfaces.

334 In practice, small changes in contact angle do not usually manifest in changes in flow properties due to the counterbalancing effects of pore structure and mineral roughness^{20,21}. The differences 335 336 in how this is manifest in trapping in the water-wet (equivalent CO₂ and N₂ trapping) and mixedwet (less CO₂ trapping) states may be due to the values of the contact angles (the pore scale force 337 balance changes as $\frac{d(\cos\theta)}{d\theta} = -\sin\theta$), or due to mineral wall roughness smoothed out by 338 339 asphaltene deposition in the mixed-wet rock. It is notable that the difference in trapping of gaseous and supercritical CO₂ observed in an unaltered carbonate in El-Maghraby and Blunt $(2012)^{11}$ also 340 341 follows the theoretical argument. Thus it is possible that the difference between those observations 342 and the observations in unaltered rock reported here are not contradictory, but rather due to a 343 difference in sensitivity of the flow properties to changes in contact angle in the respective rocks.

Finally, the decreased trapping values obtained in the altered rocks are conservative – it is likely that the irreducible residual saturation is lower. The Amott test indeed showed continued slow desaturation of the non-wetting phase with continued brine flooding of the core. It is evident that the residual trapping of CO_2 in mixed-wet carbonate reservoirs will be significantly less than in reservoir units that have not been exposed to hydrocarbon.

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350 IMPLICATIONS

On the one hand, there are a number of attributes of the use of commercial oil fields that contribute significantly to the safety and success of storage projects. The fields are well characterized. The fields also have a demonstrated fluid trap and are not as dependent on capillary and dissolution trapping as the majority of CO_2 storage sites ^{13,14}.

355 On the other hand, it is important that large scale assessments, e.g., of storage capacity take into 356 account the difference in flow physics controlling storage in oil fields relative to those of saline aquifers. A simple example using the analytic model of Juanes et al. 2010⁴⁴ illustrates the 357 358 significance of the weakened capillary trapping for CO₂ flow at the regional scale. Consider a 359 depleted oil reservoir with a thickness of 100 m with the same petrophysical properties of the Estaillades limestone used in this study - a permeability and porosity of 138 md and 0.28, 360 361 respectively. The model was used to estimate the ultimate footprint of a CO_2 plume migrating 362 horizontally below a confining caprock layer, and the timescale required for immobilization 363 through capillary trapping. Consider 100 Mt of CO₂ injected every year using 100 wells, with an interwell spacing of 1 km, for a period of 5 years. The connate water saturation and k_{ra} was 364 assumed to be 0.4 and 0.6, respectively⁴⁵. The model accounts for gravity override, capillary 365

trapping, natural groundwater flow, and the shape of the plume during the injection period but does not account for reservoir heterogeneity and changes to relative permeability due to changes in the wetting state.

In the water wet system, an initial CO₂ saturation of 42% corresponded to a 24% residual saturation. This resulted in a migration of the plume ~ 50 km prior to immobilization in ~ 5,000 years. For the same initial CO₂ saturation, only 15% was trapped in the mixed-wet system, corresponding to a migration path of ~ 130 km and immobilization in ~ 16,000 years.

Our results show that one of the key processes for maximizing CO_2 storage capacity and security is significantly weakened in many hydrocarbon reservoirs relative to saline aquifers. We anticipate this work to highlight a key issue for the early deployment of carbon storage – that those sites which are economically most appealing as initial project opportunities are the very locations in which the contribution of capillary trapping to storage security will be minimized. This should serve as a starting point for modelling studies to incorporate the reduced impact of capillary trapping on CO_2 injection projects using hydrocarbon reservoirs.

380

381 ASSOCIATED CONTENT

Supporting Information. Tables listing all IR results and corresponding 2D saturation maps of all tests are included. Derivation of the relationship between contact angles. Further information on experimental procedures, porosity and saturation measurements, a schematic of the coreflooding rig, asphaltene precipitation measurement, wetting alteration procedure, wetting state assessment, rock samples, remaining CO₂ saturations as a function of pore volume injected

387	of brine in mixed-wet system.	This material is available	free of charge via the Internet at
	5		0

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408	ABBREVIATIONS
409	CO ₂ carbon dioxide
410	N ₂ Nitrogen
411	C Land model coefficient
412	sc supercritical
413	
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