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1 Ranking the geothermal potential of radiothermal granites in 2 Scotland: are any others as hot as the Cairngorms?

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6
7 **Abstract:** Prior investigations concur that the granite plutons in Scotland which are most likely to
8 prove favourable for geothermal exploration are the Ballater, Bennachie, Cairngorm, and Mount
9 Battock plutons, all of which have heat production values greater than 5 $\mu\text{W}/\text{m}^3$. This heat
10 production arises from the significant concentrations of potassium, uranium, and thorium in some
11 granite plutons. A new field-based gamma-ray spectrometric survey targeted plutons which were
12 poorly surveyed in the past or near areas of high heat demand. This survey identifies several other
13 plutons (Ben Rhinnes, Cheviot, Hill of Fare, Lochnagar, and Monadhliath) with heat production rates
14 between 3 and 5 $\mu\text{W}/\text{m}^3$ that could well have geothermal gradients sufficient for direct heat use
15 rather than higher temperatures required for electricity generation.

16 The Criffel and Cheviot plutons are examples of Scottish granites which have concentric
17 compositional zonation and some zones have significantly higher (up to 20%) heat production rates
18 than others in the same plutons. However, the relatively small surface areas of individual high heat-
19 production zones mean that it is unlikely to be worthwhile specifically targeting them.

20 **Supplementary material:** The full set of heat production data are available at:

21 <http://dx.doi.org/10.5525/gla.researchdata.302>

22 Renewed geothermal exploration worldwide is being driven by the need to provide low carbon
23 renewable energy (Younger 2014). The last drive in British geothermal exploration concentrated on
24 identifying resources capable of generating electricity. At this time the Scottish granites were
25 identified as lacking the required geothermal resources due to the very high temperatures ($>130^\circ\text{C}$)
26 and production rates required for power generation. Renewed interest in the geothermal potential
27 of the Scottish granites has recently emerged for two reasons. Firstly, a change in proposed
28 exploitation from electricity generation to direct use heat; requiring much lower temperatures.
29 Secondly, rigorous re-assessment of heat flow leading to an upgraded geothermal resource
30 potential; suggesting higher temperatures at shallower depths than previously thought.

31 In Scotland, heat is currently responsible for over 55% of energy demand and electricity only 20%
32 (The Scottish Government 2016a). Hence the modern interest in geothermal for direct use heat
33 purposes. This paper focuses on the case of Scotland due to its acute need for renewable heat as the
34 country only generates 3% of heat from renewable sources; compared to the EU average of 16%
35 (Scottish Government 2016a). The Scottish Government has set a target of producing 11% of heat
36 from renewable sources by 2020, so to meet this target there needs to be a sudden and sustained
37 growth in renewable heat capacity. Currently, 90% of Scotland's renewable heat output is from
38 biomass (Scottish Government 2016a), however widespread uptake of biomass for district heating
39 schemes is hampered due to particulate and NO_x emission standards in cities. Geothermal resources
40 remain a hitherto untapped supply of large scale renewable heat with no emission or additional
41 infrastructure issues which could block development in cities or other locations of high heat demand
42 (e.g. distilleries and other heat-intensive industries).

43 Granite geothermal reservoirs are one of the three principal geothermal resources identified in
44 Scotland, the others being hot sedimentary aquifers and abandoned mine workings (Gillespie et al.
45 2013). Granite geothermal resources arise where plutons are naturally enriched in potassium,
46 uranium and thorium. The natural radioactive decay of these elements produces heat. Despite this
47 heat production being relatively small (on the order of several $\mu\text{W}/\text{m}^3$ of granite), the large volumes
48 of the granite plutons (e.g. Rollin 1984) means there is often sufficient heat-producing rock available
49 to materially elevate the geothermal gradient beyond that of the background heat conduction from
50 the mantle. The geothermal potential of the granites of Scotland was previously investigated as part
51 of the UK-wide assessment of the potential for new sources of electricity generation in the 1970s
52 and 80s (Downing and Gray 1986). Gamma-ray spectrometric surveys identified four high heat
53 producing granites in Scotland; Ballater, Bennachie, Cairngorm, and Mount Battock (Fig. 1). The
54 findings of these gamma-ray surveys were summarised by Downing and Gray 1986 (their Table 3.9).
55 However the original data and methods are documented in a range of sources, including in-house
56 BGS reports (Webb and Brown 1984), theses (Barritt 1983, Tindle 1982, Cassidy 1979, and
57 Hennessey 1979) as well as various reported personal communications. Different methods were also
58 reportedly used, including Instrumental Neutron Activation Analysis (INAA), X-ray fluorescence (XRF)
59 or vague “field based techniques”. Furthermore, some granite plutons only yielded small numbers
60 of heat production values, resulting in estimates which are not robust. Shallow ($\leq 300\text{m}$) heat flow
61 assessment boreholes were subsequently drilled into the four granite plutons identified as high heat
62 producing. The heat flow results from these boreholes were much lower than expected (Wheildon et
63 al. 1984) and predicted low geothermal gradients. Consequently, exploration of the Scottish granites
64 was discontinued. However, these studies neglected to correct the heat flow measurements for the
65 cooling effects from periods of lower temperatures during the Pleistocene. That past climate can
66 affect temperature in boreholes of several hundred metres depth was established decades ago (e.g.
67 Anderson 1940). This effect is expected to be particularly acute in Britain, due to the large
68 temperature differential between modern times and past ice ages largely a result of the current
69 warming effect of the gulf stream. Recent palaeoclimate corrections suggest that these granites
70 represent significantly better resources than was previously realised (Westaway and Younger 2013,
71 Busby et al. 2015). Topographical influences were also corrected for by Westaway and Younger
72 (2013) and Wheildon et al. (1984) which had the effect of reducing the heat flow estimate by
73 approximately 5%.

74 The corrected heat flow measurements combined with the current exploration being for direct use
75 heat, mean that geothermal exploration and production would be at much shallower depths than
76 previously considered. This has the effect of significantly reducing drilling costs and risk, lowering the
77 significant economic barriers to the first steps of geothermal development. Although the previous
78 gamma-ray spectrometry survey identified four high heat producing granites, other plutons were
79 poorly surveyed e.g. the Hill of Fare pluton had only two measurements. To re-evaluate the
80 radiogenic heat production properties of Scottish granite a new gamma-ray spectrometry survey has
81 been conducted. The survey aims to reduce uncertainty over granite geothermal resources in
82 Scotland by providing new heat production data for previously poorly surveyed granites. These data
83 are made freely available so any assessment using these data can better understand how the heat
84 production estimates were reached. The new data may also indicate possible granite geothermal
85 resources that have previously been overlooked.

86 Geological Setting

87 The majority of Scottish granite plutons (and those of interest in this study) were emplaced following
88 the Caledonian Orogeny and after final closure of the Iapetus during the late Silurian and Early

89 Devonian (Thirlwall 1982, Stephens and Halliday 1984, Miles et al. 2016). Biostratigraphic (Kemp
90 1987) and geochronological (Rock et al. 1986 and Kneller 1991) evidence provides a time of 420 Ma
91 for final closure of the Iapetus. The post-Caledonian granites were emplaced between 435 and 390
92 Ma (Halliday and Stephens 1984) but the majority were emplaced between 410 and 400 Ma (Soper
93 1986). Magma was likely generated due to slab breakoff during collision of Baltica with the Scoto-
94 Greenland margin (Atherton and Ghani 2002). The post-collisional regime of transpression and
95 transtension resulted in the large lateral movements along the Great Glen and Highland Boundary
96 Faults.

97 The Scottish granites inherited their isotopic signature from their host terrain (Stephens and Halliday
98 1984, Thirlwall 1989 and Canning et al. 1996) and the post-Caledonian granites were emplaced over
99 a similar time period (Miles et al. 2014). The granites are classified into four terranes: the Argyll and
100 Northern Highlands Suite, the Cairngorm Suite, the South of Scotland Suite, and the Galloway
101 Suite/Trans-Suture Suite (Highton 1999 and Miles et al. 2014). The Cairngorm Suite made up the
102 granites which were previously identified as the best geothermal granite resource and these tend to
103 have higher SiO₂ contents and have generally evolved geochemical characteristics (Downing and
104 Gray 1984). Original interpretations of the post-Caledonian granites suggested emplacement was in
105 analogous settings to Andean type subduction arcs. However there are significant compositional
106 differences between the Caledonian Granite and granite in such an Andean environment (Pitch
107 1983, Miles et al. 2014). Such disagreement testifies to the continuing research revealing ever
108 increasing emplacement complexity of the Caledonian Granite than previously thought.

109 A much younger suite of Tertiary granites in Scotland is represented in this study by the Arran
110 granite. In contrast to the Caledonian granites, the Tertiary granites formed in response to
111 emplacement of the Iceland mantle plume

112 Methods

113 Plutons Targeted

114 In Scotland, there are many more distinct granite plutons than are covered in this study. Plutons
115 were targeted in this study if they fulfilled one or more of these factors: whether the pluton had
116 relatively few existing heat production data, whether the pluton was located near places of
117 significant heat demand, and / or the pluton was previously considered a high heat producing
118 granite.

119 The Arran and Strathspey granites had no previous heat production data (locations shown on Fig. 1).
120 Many Scottish granites had fewer than ten heat production data points, these were; Aberdeen (with
121 a mean heat production of 2.2 $\mu\text{W}/\text{m}^3$, and $n=3$), Ardclach (1.4 $\mu\text{W}/\text{m}^3$, $n=4$), Ben Rhinnes (3.2
122 $\mu\text{W}/\text{m}^3$, $n=6$), Cheviot (3.0 $\mu\text{W}/\text{m}^3$, $n=6$), Grantown (1.3 $\mu\text{W}/\text{m}^3$, $n=6$), Hill of Fare (3.9 $\mu\text{W}/\text{m}^3$, $n=2$),
123 Monadhliath (5.7 $\mu\text{W}/\text{m}^3$, $n=6$), Moy (2.1 $\mu\text{W}/\text{m}^3$, $n=6$), and Ross of Mull (1.5 $\mu\text{W}/\text{m}^3$, $n=4$). Many
124 other Scottish granites also had similarly few heat production data but the surveying focussed on
125 those plutons nearer centres of heat demand and / or of larger exposed surface area. The Ballater
126 (5.7 $\mu\text{W}/\text{m}^3$, $n=34$), Bennachie (5.7 $\mu\text{W}/\text{m}^3$, $n=32$), Cairngorm (5.0 $\mu\text{W}/\text{m}^3$, $n=233$) and Mount
127 Battock (5.0 $\mu\text{W}/\text{m}^3$, $n=48$) granites were previously well surveyed but are included in this current
128 survey to allow comparison between these likely high heat producers and the granites with lower
129 heat production. Lochnagar (2.7 $\mu\text{W}/\text{m}^3$, $n=14$) and Criffel (2.2 $\mu\text{W}/\text{m}^3$, $n=14$) had fourteen and
130 eighteen measurements respectively. Despite this current low heat production estimate Lochnagar
131 is re-evaluated because of its close proximity and chemical similarity to the high heat-producing East
132 Grampian granites (Webb and Brown 1984), which suggests it may have previously been
133 underestimated. Whereas the proximity of the Criffel granite to the large heat demand of the town
134 of Dumfries justifies a re-evaluation with a more robust estimate of heat production.

135 Gamma-ray spectrometry measurements in the field

136 A GAMMA SURVEYOR II (GSII), manufactured by GF Instruments in Brno, Czech Republic, was used
137 to collect data from all the granites in this study. Using the same equipment and methods for each
138 granite ensures the readings would be consistent and directly comparable. The GSII uses a detector
139 made of Bismuth Germanate Oxide with a volume of 20cm³. The analyser measures 1024 channels
140 between 0.03 and 3 MeV. The detector material reacts with gamma-rays and produces photons of
141 visible light. These release of these photons forces electrons to be ejected from another component,
142 called the photomultiplier. These electrons then strike an anode producing a negative voltage pulse
143 proportional to the energy of the incident gamma-ray photon. Analyzers within the GSII then use the
144 combined pulses over minutes of readings to produce estimates for potassium, uranium, and
145 thorium contents as radioactive decay from these elements produces gamma-rays of different
146 energies. For further details of the workings of gamma-ray spectrometry and geothermal
147 exploration we refer the reader to McCay et al (2014). The GSII was chosen for these surveys
148 because it is not prone to significant signal drift caused by temperature and other influences, and
149 thus does not require daily calibration. It is also a compact and relatively light device (weighing only
150 a few kilograms) so is easy to carry in the mountainous and remote terrains which host most of the
151 granite exposure in Scotland. The output of the GSII provides information on the concentrations of
152 potassium (as percentage by weight), uranium and thorium (both as parts per million). These
153 concentrations are readily converted into equivalent radiogenic heat-production rates, as explained
154 below.

156 To take measurements the GSII was placed directly against an exposed area of granite. Gamma-rays
157 can penetrate up to 0.5 m through rock, however the half-lengths of gamma-rays mean that the
158 majority of the detected photons will come from the shallowest 15cm of rock with a smaller
159 contribution from deeper sources. This effectively results in the sample collecting information from a
160 15cm thick disc or rock which is 1m in radius (McCay et al. 2014). Raising the GSII above the ground
161 would allow a much larger sample size to be effectively measured (e.g. Løvborg et al. 1979), due to
162 gamma-rays being able to penetrate several hundred metres through air. However this elevated
163 reading would be affected by the soil and vegetation cover on the rock so we opted for direct
164 measurements to get a less ambiguous representation of granite heat production rates. The GSII is
165 calibrated to take readings from a flat area of rock. If the measurement is taken in a depression then
166 the readings will be overestimated due to gamma-rays from the overlooking areas of rock hitting the
167 spectrometer, whereas if the measurement is on a mound that the readings will be underestimated
168 (McCay et al. 2014). Measurement time was between 3 to 5 minutes as this gave the best balance
169 between accuracy and speed. For each sample, at least two measurements were taken, if these
170 were not sufficiently similar (within 10% of the U value) then further readings were taken until
171 consistent readings were gathered.

173 Data Processing

174 The concentrations of potassium, uranium, and thorium provided by the GSII were used in the
175 following equation to estimate heat production (HP):

$$HP(\mu W m^{-3}) = \rho (0.035C_K + 0.097C_U + 0.026C_{Th}) \quad (1)$$

176 Where: ρ is rock density (kg m⁻³), C_K is concentration of potassium by % weight, C_U and C_{Th} are
177 concentration of uranium and thorium in ppm. The potassium % weight used for heat production is
178 elemental, as opposed to percentage weight of K₂O used in other fields such as hydrogeology. The
179 constants before each elemental concentration in equation 1 are based on the energy released
180 during alpha, beta, and gamma decay of the radioelements (Birch 1954 and Rybach 1976).

181 Given the practical difficulty in obtaining and transporting granite samples, and the protected status
182 of some of the sites we surveyed, we did not take samples from each measurement location in order
183 to find the density of granite. Instead we assumed a density of granite of 2.7g/cm^3 for all samples
184 based on information from the (Rollin 2009). For an example of how this assumption may influence
185 the results, a sample with heat production of $4\mu\text{W/m}^3$ with a density of 2.6 would have a change of
186 heat production of $\pm 0.2 \mu\text{W/m}^3$ with a variation in density of $\pm 0.1 \text{kg/m}^3$. Although this is not an
187 insignificant influence on heat production it is not enough for a granite pluton to be incorrectly
188 identified (or dismissed) as of high heat production.

189 The arithmetic mean of the measurements was used to create an estimate of heat production for
190 each pluton. Geometric means might be more appropriate in such a geological distribution due to
191 the arithmetic mean losing information about the small scale heterogeneities in heat production
192 observed in some locations, which could be due to the incremental emplacements of plutons over
193 tens of millions of years (e.g. Glazner et al. 2004). For example, the Mount Battock pluton has an
194 area approximately 2m^2 where heat production is $15\mu\text{W/m}^3$. However arithmetic mean is preferred
195 to be consistent with the previous literature on heat production data in the United Kingdom (e.g.
196 Downing and Gray 1986). Any measurements which were noted in the field as being influenced by
197 nearfield topography (Tyler 2000) were removed prior to averaging (for information these data are
198 still included and clearly identified in the database linked to this paper). Many of the measurements
199 from the Hill of Fare granite were overestimates because they were collected within abandoned
200 quarries. A bespoke method was developed to deal with these issues in this plutons which is
201 explained below. Finally the calculated arithmetic means were combined with those of previous
202 studies presented in Downing and Gray (1986).

203

204 Correction for overestimation in Quarry Measurements

205 Many sample locations of the Hill of Fare pluton were taken from abandoned quarries. These
206 quarries have the advantage of being relatively fresh faces as they have not exposed to erosion for
207 as long as the natural outcrops. However 10 of these sample locations had the disadvantage that the
208 nearby quarry walls and ledges would lead to overestimated results. There are no standard methods
209 detailed in the gamma-ray spectrometry literature to deal with such problems, so a novel solution is
210 proposed here. Potassium shows the most consistent concentrations of the three radio-elements in
211 granites so it is the variation in uranium and thorium concentrations that generally determine higher
212 and lower heat production zones. Weight percent K in granite is typically between 3.8 and 4.5 for
213 the East Grampian Siluro-Devonian granites (Table 3.9 Downing and Gray 1986), which is the suite to
214 which the Hill of Fare granite belongs. However values of potassium concentration in the quarries
215 were consistently higher than 5.0 indicating that they are likely overestimates due to the other
216 quarry walls and surfaces.

217 The GSII is calibrated to be accurate when placed on a flat surface, the spectrometer effectively
218 samples radiation incident from a solid angle of 2π steradians or a half space (Fig. 2). Hypothetically,
219 a spectrometer placed underground would sample from an angle of 4π steradians or a full space, as
220 it would be completely surrounded by radiogenic rock (Fix x – quarry effective sample volume). The
221 shape of the quarries meant that the sample was effectively from 3π steradians or a three quarters
222 space. This 3π sample volume should lead to an overestimation of the concentrations of
223 radioelements by approximately 50%.

224 Such an overestimation would affect all three radio-element estimations by the same proportion. So
225 if a “true” value of one of the radio-elements was known, then the ratio between that true value and
226 the spectrometer estimation would indicate the degree of overestimation. The concentrations of
227 potassium have previously been relatively consistent in high heat production granites compared

228 with uranium and thorium. Downing and Gray (1984) report that high heat producing granites in
229 Scotland have average potassium concentrations ranging between 4.1% and 4.4%. On the Hill of Fare
230 granite, flat surface measurements were consistent with this trend giving values ranging between
231 4.0% and 4.8%. When tested against quarry faces (i.e. the quarry effective sample volume on Fig. 2)
232 the GSII gave a potassium concentration of 7.14%. Which is 60-70% higher than what would be
233 expected from a flat surface, but in-line with the enhancement expected from a 3π steradian, or
234 greater, sample volume.

235 From the above test a method was developed to normalise potassium, uranium, and thorium
236 values. A “true” value of 4.5 was assumed for potassium concentrations. Then, all three of the
237 radioelements were divided by the ratio between the measured potassium and the assumed
238 “correct” value of 4.5. For example, if a potassium concentration value in the quarries was measured
239 as 9.0, then the correction factor is $4.5/9.0$ which is 0.5. The potassium, uranium, and thorium
240 concentrations would need to be halved to achieve a corrected estimate of their concentrations. In
241 total, ten sample locations underwent this correction for quarry topography in the Hill of Fare
242 granite.

243 Results

244 Heat production values of granite plutons

245 New gamma-ray spectrometry data were collected from 18 granite plutons (Table 1 and Fig. 3). The
246 granites show a range of heat production values, with Strathspey the lowest at $1.20 \mu\text{W}/\text{m}^3$ and
247 Ballater the highest at $8.22 \mu\text{W}/\text{m}^3$. Generally, with heat production any granite greater than 4
248 $\mu\text{W}/\text{m}^3$ is of interest for direct-use geothermal exploration as this value tends to be high enough to
249 significantly raise the geothermal gradient over that of the surrounding rock (Gillespie et al. 2013).
250 Additionally, $4 \mu\text{W}/\text{m}^3$ corresponds to approximately a value of 10 in the old imperial units for heat
251 production (i.e. 10^{-13} calories/cm³). Naturally this is a guide and dependent upon local conditions; for
252 instance, the unexposed Weardale granite in County Durham, England, has a heat production of 4.11
253 $\mu\text{W}/\text{m}^3$ but its insulating Carboniferous sedimentary overburden allows much higher temperatures
254 to build up than the granite was exposed at the surface (Manning et al. 2007).

255 For this paper we adopt the convention that any granite pluton with a mean surface heat production
256 above $5 \mu\text{W}/\text{m}^3$ is considered high heat production, between 3 and $5 \mu\text{W}/\text{m}^3$ is considered
257 marginally high heat production (referred to as marginal heat production), and below $3 \mu\text{W}/\text{m}^3$ is low
258 heat production. Four of the surveyed granites have high values of heat production; Ballater,
259 Bennachie, Cairngorm, and Mount Battock (shown on Fig. 1). Five of the granites have marginal
260 values of heat production; Ben Rhinnes, Cheviot, Hill of Fare, Lochnagar, and Monadhliath (shown in
261 yellow on Fig. 1). Eight of the surveyed granites have low heat production; Aberdeen, Ardclach,
262 Arran, Criffel, Strathspey, Grantown, Moy, and Ross of Mull (shown as grey on Fig. 1). The past heat
263 production estimates were classified as “satisfactory” if there was sufficient quantity and spatial
264 distribution of data (Downing and Gray 1986), however it was not stated what this meant. However,
265 it is possible to infer that those plutons previously labelled as having satisfactory heat production
266 estimates had more than ten samples from several locations in a pluton; for consistency this
267 classification will be used in the following discussion.

268 The Lochnagar granite can be classified as marginal heat production, with a value of $3.89 \mu\text{W}/\text{m}^3$.
269 The lower historical value ($2.7 \mu\text{W}/\text{m}^3$) was classified as low confidence, possibly due to only having
270 14 heat production samples from the Lochnagar pluton. Downing and Gray (1984) classify granites
271 with similarly few data as robust estimates, but do not specify why Lochnagar is not considered a
272 robust estimate.

273 The Hill of Fare pluton previously only had two data points, far too few for the estimate of 3.9
274 $\mu\text{W}/\text{m}^3$ to be considered representative of the pluton. The new estimate, consisting of 31 data
275 points, give a surface heat production average for of $\mu\text{W}/\text{m}^3$. The Hill of Fare therefore classified as a
276 marginal heat producer.

277 The heat production estimate for the Monadhliath pluton is $4.98 \mu\text{W}/\text{m}^3$. Although this is a marginal
278 classification, it is only $0.02 \mu\text{W}/\text{m}^3$ below a high heat production classification, it could therefore be
279 reasonably classified as a marginal/high heat production granite. Only six measurement informed
280 the previous heat production estimate for the Monadhliath pluton, but with the addition of the 17
281 new measurements then the current estimate would typically be considered representative of the
282 pluton (i.e. following the method of Downing and Gray 1984). Lochnagar, Hill of Fare and
283 Monadhliath are located relatively close to each other (Fig. 1) and are part of the same geochemical
284 suite as the four high heat producing granites (Busby et al. 2015).

285 Cheviot and Ben Rhinnes were both previously estimated as low heat producers but these did not
286 have enough data for confidence in these estimations (Fig. 4 and Table 1). The new heat production
287 data indicate the Cheviot granite is marginal at $3.70 \mu\text{W}/\text{m}^3$ which is higher than the previous
288 estimate of $3.00 \mu\text{W}/\text{m}^3$. There are still few data available for the Ben Rhinnes pluton, due to lack of
289 exposure, but a combined mean of $3.33 \mu\text{W}/\text{m}^3$ suggests this granite is likely to be at the lower end
290 of the marginal category.

291 Of the eight plutons classified as low heat production only Arran and Criffel previously had
292 significant number of measurements, at 20 and 93 respectively. The plutons Moy, Ross of Mull, and
293 Grantown had over ten measurements when the new and old data were combined. Even with the
294 combination of new and old data Aberdeen (six measurements), Ardclach (seven measurements),
295 and Strathspey (five measurements) are still currently poorly surveyed. The lack of exposure of these
296 granites inhibits further field campaigns adding to these totals. However no measurements of these
297 plutons were much greater than the mean, as shown by the small error bars on figure 3, which does
298 not suggest that further measurements would be likely to reclassify these plutons.

299 Zoned plutons

300 Some granite plutons are concentrically zoned, with the zones differentiated on the basis of
301 geochemical/mineralogical changes in composition (e.g. Stephens et al. 1985). Here we investigate
302 whether these individual zones represent distinctive areas of high heat production within a pluton.
303 Sufficient data were collected from two zoned plutons which would allow robust comparison
304 between the individual zones. These are the Criffel Pluton in near Dumfries in South West Scotland,
305 and the Cheviot Pluton which lies on the eastern area of the Scottish-English border (see Fig. 1 for
306 specific locations).

307 *Criffel Pluton*

308 Work by Stephens et al. (1985), Stephens (1992) and continued by Miles et al. (2013) identified five
309 zones of granite forming concentric rings within the Criffel Pluton (Fig. 5). The outer two zones (1
310 and 2) are granodiorites while the inner three zones (3, 4 and 5) are granites. The general trend is for
311 the granite to become more silicic towards the inner zones, with the outer zone having SiO_2 at
312 approximately 65% compared with 72% of the inner most zone (Stephens and Halliday 1980).

313 The highest value of heat production is in zone 2 (the most outer zone but one) with $2.51 \mu\text{W}/\text{m}^3$,
314 and the lowest is in the central zone 5 with $1.83 \mu\text{W}/\text{m}^3$ (Fig. 6 and Table 2). From zone 2 to zone 5
315 there is a general trend for heat production to decrease towards the centre of the pluton. However
316 the outermost zone 1 does not follow this trend. Although there are clear differences in the heat

317 production, none of the zones are above $3 \mu\text{W}/\text{m}^3$. All zones are considered as low heat production
318 and unlikely to host enhanced geothermal gradients due to the natural decay of radioactive
319 material.

320 *Cheviot Pluton*

321 The first detailed geological mapping of the Cheviot pluton was conducted by Jhingran (1942), who
322 identified three main types of granite, an outer Marginal zone (1), an middle Granophyric zone (2),
323 and an inner Standrop zone (3); with a fourth zone being an area that displays properties of the
324 outer zones 2 and 1 (Fig. 7). The Cheviot pluton is surrounded by contemporaneous lavas. Al-Hadif
325 (1985) subsequently produced a more detailed map of the Cheviot pluton; however this has not
326 been peer-reviewed and published. Additionally we do not have enough data for the Cheviot pluton
327 to provide robust differentiation between many granite types, for these reasons we use the Jhingran
328 (1942) classification. The whole pluton is classified as granodiorite, except for some small areas of
329 the middle Granophyric zone 2 which are true granite.

330 The outer marginal Zone 1 has the highest heat production of $4.16 \mu\text{W}/\text{m}^3$, which would be enough
331 to classify it as a marginal heat producing granite (Fig. 8 and Table 3). Zone 3 has a lower heat
332 production of $3.51 \mu\text{W}/\text{m}^3$, zone 2 also appears relatively low however only having one
333 measurement means that this value cannot be considered representative of the zone. The
334 surrounding lavas heat production, of $3.43 \mu\text{W}/\text{m}^3$ is lower than that of the granites, however this
335 value is surprisingly high for andesites. A possible explanation for the high heat production is the
336 hydrothermal alteration of feldspar to saussurite (Lee 1982, Al-Hafdh 1985). Hydrothermal
337 Alteration breaks down the structure of the minerals allowing remobilisation of the radioelements,
338 which can result in high concentrations of radio-elements away from their emplacement location.

339 Discussion

340 Are there any reclassifications of geothermal potential for any Scottish granites?

341 Figure 4 compares our new heat production estimates with previous estimates for those plutons (cf
342 Downing and Gray 1986 - their Table 3.9). Any significant divergence from the equivalence line
343 would indicate a pluton that may require re-evaluation. Table 1 shows the mean surface estimate
344 values plotted on figure 4 and then number of data used for each mean, in addition to the mean
345 surface heat production of the combined new and historical data.

346 The four East Grampian granite plutons of Ballater, Bennachie, Cairngorm, and Mount Battock all
347 have heat production rates in excess of $5 \mu\text{W}/\text{m}^3$, and are robustly characterised as high heat
348 producers by both the new and historical data (Fig. 4 and Table 1). All four of these plutons are
349 classified as part of the Cairngorm suite of post-Caledonian granites (Highton 1999). The new and
350 historical data agree that Cairngorm, Mount Battock, Ballater, and Bennachie plutons are all high
351 heat producers. However, the new estimates are generally higher than the historical estimates but
352 not to the degree to require a significant re-appraisal of their geothermal potential. The new data
353 show that four granite plutons should be reconsidered potential resources as marginally high: Hill of
354 Fare, Monadhliath, Lochnagar, and Cheviot. The Hill of Fare, Lochnagar, and Cheviot granites were
355 previously overlooked due to lack of data, and the few data available appeared to indicate that they
356 would be low heat producers. Monadhliath was previously considered a high heat producer, but the
357 few heat production data and poor accessibility meant it has achieved less attention than the four
358 other high heat production granites. All of these plutons, with the exception of the Cheviot, are part
359 of the Cairngorm suite, the same as the high heat producing granites..

360 The Hill of Fare pluton had only two previous measurements of heat production which gave a mean
361 value of $3.9 \mu\text{W}/\text{m}^3$ (Table 1). The new data ($n=29$) give a more robust value of $4.03 \mu\text{W}/\text{m}^3$ when
362 combined with the two previous data, which classifies the granite as a marginal heat producer.

363 The Monadhliath granite was previously thought likely to be a high heat producer and potential
364 geothermal resource. However, only six heat production data were collected on the Monadhliath
365 granite. The new 17 data points show the Monadhliath granite being a comparable resource to the
366 four high heat producing granites with a heat production of $4.98 \mu\text{W}/\text{m}^3$, which is barely under the
367 threshold to be considered a high heat producer.

368 The Lochnagar and Cheviot granites have, hitherto, been disregarded as potential geothermal
369 resources, as they apparently had low heat production values of $2.7 \mu\text{W}/\text{m}^3$ and $3.0 \mu\text{W}/\text{m}^3$
370 respectively (Table 1). However, when our new data are considered then the Lochnagar granite is
371 classified as a marginal heat producer at $3.89 \mu\text{W}/\text{m}^3$ and the Cheviot granite is only slightly lower at
372 $3.70 \mu\text{W}/\text{m}^3$.

373 What is the significance of classifications for geothermal exploration?

374 The new gamma-ray spectrometry results confirm the classification of the high heat production
375 granites i.e. Bennachie, Ballater, Cairngorm, and Mount Battock. These four granites are the most
376 likely to have useful temperatures for direct use heat exploitation at the shallowest depths of the
377 Scottish granites. And would therefore, present the most promising resources if there were nearby
378 heat demands. This study also identifies those granites with low heat production, which are not
379 expected to have useful temperatures at depths for an economic geothermal scheme. These low
380 heat production granites include Arran and Strathspey on which there were previously no heat
381 production data. All, apart from Grantown, of the low heat production granites previously had few
382 measurements leading to uncertainty in their classification (Table 1). An example of the poor heat
383 resource of these granites comes from a borehole (NJ91SE3 at UK national grid reference 395090 810900)
384 which was drilled into country rock 3km adjacent to the Aberdeen granite. The borehole recorded a
385 temperature of 32°C at 1494m (Groves et al. 2012). Although the borehole is not directly in the
386 granite a higher geothermal gradient would be expected if the granite itself had significantly
387 elevated temperatures. The geothermal gradient found in this borehole is very low at around 13.4°C
388 per km. This would require such great depths to be drilled to reach useful temperatures, that even a
389 direct use geothermal project would be unlikely to be economic. Removing the uncertainty that
390 these granites have poor geothermal resource, means that exploration effort can be focussed on
391 those granites with favourable resources potential.

392 More crucially, marginal heat production granites have been identified, these are the Lochnagar,
393 Monadhliath, Hill of Fare, and Cheviot granite as well as the Ben Rhines pluton to a lesser extent.
394 These plutons could be viable geothermal resources if a heat-demanding development be proposed
395 within close proximity to these granites. The Hill of Fare Pluton is of particular interest because it lies
396 close to the large heat demand of the growing Banchory District Heating Scheme. For this reason, a
397 feasibility study investigated using the Hill of Fare granite as a geothermal resource for the Banchory
398 District Heating Scheme (The Scottish Government 2016b). The study predicted temperatures at
399 depth within the Hill of Fare Pluton (Table 4). These predicted geotherms are sufficient for direct use
400 geothermal to compete economically with a natural gas boiler over the lifespan of a project; in
401 addition to the significant emissions advantage over natural gas and air quality advantages over
402 biomass. This suggests, that geothermal exploration in Scotland can rightly expand, beyond merely
403 the four high heat production granites, to also consider the marginal heat production granites as
404 targets.

405 Robust values of heat production are the first step in the geothermal exploration process of
406 radiothermal granites. The simple ranking system, proposed in this paper, builds the framework
407 upon which the next stage of the exploration process can be targeted towards the granites which
408 are most likely to have favourable geothermal gradients, i.e. the high and marginal heat production
409 granites. The next stage in the appraisal of these granites would require deep scientific boreholes of
410 1000m and greater depth. The kilometre depth of these scientific boreholes is required to provide
411 certainty of the heat resource, due to past experience of the misleading heat flow that shallow
412 boreholes can present (Westaway and Younger 2013 and Busby et al. 2015). The scientific borehole
413 would provide information about heat and permeability at depth (cf Manning *et al.* 2007; Younger
414 and Manning 2010), which would create the information required to justify commercial exploration
415 boreholes.

416 Could the higher heat production outer rims of the zoned granites be a potential
417 geothermal resource?

418 Significant variations in heat production can exist with a pluton made up of distinct geochemical
419 zones. Here, we discuss the possibility of these distinct high heat production zones offering a
420 potential geothermal exploration target within an otherwise discounted granite. We do not explore
421 the emplacement and post-emplacement processes which have led to zonal differences in
422 concentrations of potassium, uranium, or thorium but focus on the potential geothermal resource
423 that these differences could pose. However there is much scientific and applied merit in
424 investigating these processes further as they could elucidate how heat production may change with
425 depth (particularly in relation to 3D geometry) and help reduce the risk of the geothermal
426 exploration process within granite.

427 The investigation of the Criffel pluton shows that there were heat production differences between
428 the geochemically distinct zones of the granite (Figures 7 and 8). However none of these zones were
429 of high heat production, so these differences are not considered significant for geothermal
430 exploration of this pluton.

431 In the Cheviot granite all the zones would also individually be classed as a marginal heat producer,
432 except for the outermost zone 1 which has a heat production of about $4 \mu\text{W}/\text{m}^3$. Zone 1 was
433 significantly higher (by $0.47 \mu\text{W}/\text{m}^3$) than the next highest heat production which was zone 2,
434 however zone 2 only had one measurement so can not be considered a reliable estimate. This outer
435 zone 1 may have high enough heat production to be regarded as a possible geothermal target but it
436 has a relatively small surface area compared to the uniformly high heat producing granites. For
437 comparison the surface area of the Bennachie pluton is 55 km^2 whereas the surface area of the
438 marginal zone 1 of the Cheviot Pluton is 14 km^2 . Zone 1 only represents 25% of the entire Cheviot
439 Pluton surface area. There are some Scottish granites which are high heat producing and have small
440 surface areas, e.g. the Fearn Pluton ($5.1 \mu\text{W}/\text{m}^3$, 12 km^2) located north of the Great Glen Fault (Fig.
441 3.3 in Downing and Gray 1986). These small surface area granites have not been considered for
442 geothermal resources as their likely low volume at depth may not allow heat to sufficiently
443 accumulate from the heat production. Although their surface area may be misleading and actually
444 be part of a significant volume of granite at depth (i.e. similar to the tip of the iceberg). However the
445 small surface area of high heat production couple with the uncertainty of the three dimensional
446 shape of the granites mean that these zoned granites are unlikely to be favourable exploration
447 targets. For these reasons, exploration has been focussed on those granites known to have
448 significant subsurface volumes due to their large surface areas. The heat production data do not
449 support a change to this approach. For the Cheviot Pluton a model of how the different granite types
450 are distributed at depth would help constrain how much of the volume of the Cheviot Pluton is the

451 higher heat producing zone. The Cheviot Pluton could still represent an interesting geothermal
452 target if it is likely that zone 1 encompasses a larger proportion of the volume than its surface area
453 suggests.

454 Conclusions

455 High heat production Scottish granites are identified as the Ballater, Bennachie, Cairngorm, and
456 Mount Battock plutons. These four are likely to be the best granite-based geothermal resources in
457 Scotland.

458 Marginal heat production granites are identified as the Ben Rhinnes, Cheviot, Hill of Fare, Lochnagar,
459 and Monadhliath plutons. Although these granites may not present such a favourable geothermal
460 resource as the high heat producers, they are still worth investigating as a supply to any proximal
461 large heat demand.

462 Low heat production granites are identified as Aberdeen, Ardclach, Arran, Criffel, Strathspey,
463 Grantown, Moy, and Ross of Mull. These granites are very unlikely to offer geothermal gradients of
464 interest for economical direct use of heat. As such, they would be unlikely to be worth further
465 exploration if a large heat demand existed in the vicinity of these granites.

466 Zoned granites are unlikely to present geothermal resources within specific geochemical zones. The
467 relatively small surface area and likely small volume of any single higher heat production zone means
468 that they are unlikely to create large reservoirs of elevated temperature at depth.

469 This study shows the value that a relatively quick and simple gamma-ray spectrometry survey can set
470 initial targets for geothermal exploration into granite plutons. The classifications of high, marginal,
471 and low heat production provide a basis for exploring the possibility of exploiting direct-use
472 geothermal heat by any future development with large heat demands.

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631

632 Figure Captions

633

634 **Fig. 1.** Map showing locations of major exposed granite plutons in Scotland. Numbered plutons are
635 those surveyed in this study. 1 Moy, 2 Ardclach, 3 Ben Rhinnes, 4 Grantown, 5 Monadhliath, 6
636 Bennachie, 7 Cairngorm, 8 Ballater, 9 Hill of Fare, 10 Aberdeen, 11 Strathspey, 12 Lochnagar, 13
637 Mount Battock, 14 Ross of Mull, 15 Arran, 16 Cheviot, 17 Criffel. Plutons identified in this study as
638 high heat producing are 6, 7, 8, and 13, those identified as marginally high heat producing are (3, 5,
639 9, 12, and 16).

640 **Fig. 3.** Schematic of simplified quarry topography. The GSII spectrometer is calibrated assuming a 2π
641 sample volume, but the quarry topography led to an effective sample volume of approximately 3π .

642 **Fig. 3.** Mean surface heat production estimates for surveyed plutons in this study. Whiskers on the
643 icons are the standard error of the mean (standard deviation divided by the square root of the
644 number of data).

645 **Fig. 4.** Mean surface heat production estimates from figure 3 (y-axis) compared with historical
646 estimates (x-axis) given in table 3.9 in Downing and Gray (1986). Dotted line shows equivalence line.
647 Hollow circles show the estimates which Downing and Gray (1986) stated there was inadequate data
648 to derive a satisfactory heat production estimate. Arran and Strathspey granites are not included in
649 this graph as these plutons had no previous heat production data.

650 **Fig. 5.** Map of the different geochemical zones within the Criffel Pluton, after Miles et al. (2013).
651 Dotted lines show areas where sample points are clustered.

652 **Fig. 6.** Mean surface heat production values for the zones of the Criffel Pluton .

653 **Fig. 7.** Map of the different geochemical zones within the Cheviot Pluton, after Jhingran (1942).
654 Dotted lines show areas where sample points are clustered.

655 **Fig. 8.** Mean surface heat production values for the zones of the Cheviot Pluton. Note zone 2 only
656 had one measurement so does not include standard error.

657 **Table 1** *Number of data and surface heat production estimate for each pluton from the new data*
658 *presented in this paper and the historical data from Downing and Gray (1986; their Table 3.9), and*
659 *then the combined surface heat production estimate from new and historical data. *Indicates*
660 *historical data with low confidence attached to the estimate. + criteria for re-evaluation are classified*
661 *as follows: (1) few existing heat production data, (2) located near places of significant heat demand,*
662 *(3) previously considered a high heat producing granite*

663 **Table 2** *Mean surface heat production values for the zones of the Criffel Pluton.*

664 **Table 3** *Mean surface heat production values for the zones of the Cheviot Pluton*

665 **TABLE 4** *Predicted temperatures in the Hill of Fare granite for three "heat scenarios" from The*
666 *Scottish Government (2016b). Temperatures were calculated assuming a related but lower heat flow*
667 *to that measured in the nearby high heat producing granites.*