



Sampson, C., Smith, A., Bates, P. D., Neal, J. C., & Trigg, M. A. (2015). Perspectives on open access high resolution digital elevation models to produce global flood hazard layers. *Frontiers in Earth Science*, 3(85), [00085]. DOI: 10.3389/feart.2015.00085

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Perspectives on open access high resolution digital elevation models to produce global flood hazard layers

Christopher C. Sampson^{1*}, Andrew M. Smith¹, Paul D. Bates¹, Jeffrey C. Neal¹, Mark A. Trigg¹

¹School of Geographical Sciences, University of Bristol, United Kingdom

Submitted to Journal:
Frontiers in Earth Science

Specialty Section:
Hydrosphere

ISSN:
2296-6463

Article type:
Perspective Article

Received on:
27 Aug 2015

Accepted on:
02 Dec 2015

Provisional PDF published on:
02 Dec 2015

Frontiers website link:
www.frontiersin.org

Citation:
Sampson CC, Smith AM, Bates PD, Neal JC and Trigg MA(2015) Perspectives on open access high resolution digital elevation models to produce global flood hazard layers. *Front. Earth Sci.* 3:85. doi:10.3389/feart.2015.00085

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1 **Perspectives on open access high resolution digital elevation models to produce global**
2 **flood hazard layers**

3 Christopher C. Sampson¹, Andrew M. Smith², Paul D. Bates², Jeffrey C. Neal², Mark A. Trigg²

4 ¹*Corresponding author, School of Geographical Sciences, University of Bristol, BS8 1SS, UK,*
5 *chris.sampson@bristol.ac.uk*

6 ²*School of Geographical Sciences, University of Bristol, BS8 1SS, UK*

7

8 **Abstract**

9 Global flood hazard models have recently become a reality thanks to the release of open access global
10 digital elevation models, the development of simplified and highly efficient flow algorithms, and the
11 steady increase in computational power. In this commentary we argue that although the availability of
12 open access global terrain data has been critical in enabling the development of such models, the
13 relatively poor resolution and precision of these data now limit significantly our ability to estimate flood
14 inundation and risk for the majority of the planet's surface. The difficulty of deriving an accurate 'bare-
15 earth' terrain model due to the interaction of vegetation and urban structures with the satellite-based
16 remote sensors means that global terrain data are often poorest in the areas where people, property (and
17 thus vulnerability) are most concentrated. Furthermore, the current generation of open access global
18 terrain models are over a decade old and many large floodplains, particularly those in developing
19 countries, have undergone significant change in this time. There is therefore a pressing need for a new
20 generation of high resolution and high vertical precision open access global digital elevation models to
21 allow significantly improved global flood hazard models to be developed.

22

23 **Article**

24 Around the turn of the millennium, high quality two dimensional hydraulic models capable of
25 simulating the dynamics of flood inundation became a reality at the reach scale as a result of faster
26 computers, improved algorithms (Bates & De Roo 2000; Bradford & Sanders 2002; Bradbrook et al.

27 2004) and new forms of rapidly-collected remotely sensed digital elevation models (DEMs) (Bates
28 2004; Marks & Bates 2000; Bates et al. 2003; Cobby et al. 2001; Sanders 2007). Of particular value
29 to hydraulic modellers in developed countries was the commencement of routine LIDAR collection
30 due to its high horizontal and vertical precision and accuracy, its ability to penetrate vegetation cover
31 and its reduced susceptibility to scatter and shadowing relative to other forms of remotely sensed
32 elevation data such as Interferometric Synthetic Aperture Radar (InSAR) (Bates 2004). These three
33 key properties made it ideally suited to the creation of 'bare-earth' Digital Terrain Models (DTMs), a
34 type of DEM in which surface features such as vegetation and built structures are removed to leave, as
35 the name suggests, a three dimensional representation of the bare-earth surface. Such data are ideally
36 suited for the purposes of flood hazard simulation using hydraulic models, and form the basic datasets
37 from which developed world flood hazard layers, such as the Federal Emergency Management
38 Agency (FEMA) flood maps in the USA, and the Environment Agency Flood Maps in the UK, are
39 produced.

40 Whilst there have been significant advances in the models and data available for relatively small scale
41 modelling of flood inundation where high quality terrain data exist, the computational and data costs
42 associated with such models tends to restrict their application to populated areas in wealthier nations.
43 Furthermore, due to the potential impact on property prices and local economies, local or national
44 authorities may be reluctant to release the results of such models even where they do exist. However,
45 flood risk is very clearly a global problem and, consequentially, a number of research and commercial
46 groups are currently working on the development of flood hazard models at the global scale (Ward et
47 al. 2015; Winsemius et al. 2013; Hirabayashi et al. 2013; Hallegatte et al. 2013; Sampson et al. 2015).
48 Projections of rapidly escalating economic losses due to flooding (Hallegatte et al. 2013) and the
49 United Nations (UN) adoption of both the Sendai Framework for Disaster Risk Reduction (United
50 Nations General Assembly 2015) and the Warsaw International Mechanism for Loss and Damage
51 Associated with Climate Change Impacts (United Nations Framework Convention on Climate Change
52 2013) provide clear motivation for the development of global flood risk assessments by both
53 commercial and governmental purposes. Much like the earlier development of reach scale models,

54 these new models are becoming tractable as a result of further increases in computational power and
55 software parallelisation (Neal et al. 2010; Lamb et al. 2009), algorithmic improvements (Bates et al.
56 2010), and emerging global datasets (Elvidge et al. 2007; Lehner et al. 2008; Smith et al. 2015; Jarvis
57 et al. 2008; Andreadis et al. 2013; Yamazaki et al. 2014). The data challenges are particularly
58 onerous because, whereas at the reach scale most of the required ‘secondary’ spatial data other than
59 the DEM (such as river locations, channel geometries and flood defences) can viably be obtained
60 using manual survey or are contained in the data produced by national mapping agencies, at the global
61 scale all such data must be derived in an automated or semi-automated manner from remotely sensed
62 data. The DEM is the core dataset from which many of these secondary datasets are derived and, as
63 we argue in this perspective, it is the limited quality of the present generation of global DEMs that
64 presents the greatest challenge to flood inundation modellers today.

65 Although a number of free and commercial global DEMs exist, two in particular have received the
66 majority of attention from flood modellers: the Shuttle Radar Topography Mission (SRTM) (Rabus et
67 al. 2003; Farr et al. 2007) DEM and the Advanced Spaceborne Thermal Emission and Reflection
68 Radiometer (ASTER) (Abrams 2000) DEM, and their respective derivatives (Jarvis et al. 2008;
69 Kobrick 2013; Fujisada et al. 2012). These data sets are popular because they are open access and
70 offer greater levels of detail than the previous generation of open access DEMs (such as ACE GDEM
71 (Berry et al. 2000), GLOBE and GTOPO30) due to their greatly increased resolutions. For example,
72 ASTER and SRTM have ground spatial resolutions of 1 arc-seconds (~30 m at the equator
73 respectively, compared to ~30 arc-seconds (~1 km) for the previous generation DEMs. A number of
74 studies (Hirt et al. 2010; Rexer & Hirt 2014; Jing et al. 2013; Jarihani et al. 2015) have compared the
75 SRTM and ASTER DEMs across a range of locations globally to assess their applicability to
76 hydraulic models (e.g. Sanders 2007), and despite its lower nominal resolution it is SRTM –
77 particularly the void-filled CGIAR-CSI version 4 variant (Jarvis et al. 2008) - that has emerged as the
78 favoured choice. This is due to SRTM’s greater feature resolution, reduced number of artefacts and
79 lower noise than ASTER, particularly in the flatter areas of concern to flood modellers (Jing et al.
80 2013; Rexer & Hirt 2014). The prohibitive cost and restricted rights associated with commercial

81 DEMs (such as the Intermap Nextmap[®] World 10[™] and World 30[™], and Airbus WorldDEM[™], data
82 sets) restricts significantly the application of such products. This results in limited (or no) public and
83 independent validation of commercial DEMs, a lack of independent studies comparing them to other
84 DEMs, and a lack of the types of derived datasets, such as global hydrography data, that have
85 emerged from their open access counter-parts.

86 User generated 'secondary' datasets derived from global topography offer a valuable resource for a
87 range of activities and can be directly attributed to the production of open access global DEMs.
88 From a flood modelling perspective perhaps the most valuable example is the Hydrosheds global
89 hydrography dataset (Lehner et al. 2008). This dataset was produced by executing a number of
90 hydrology-based GIS operations over a suitably void-filled SRTM dataset, and contains layers such as
91 flow direction maps, river networks (with upstream accumulation areas) and catchment masks. The
92 Hydrosheds data has been used as the basis for a number of large scale hydrology and river routing
93 models (Wood et al. 2011; Yamazaki et al. 2013; Sampson et al. 2015; Gong et al. 2011; Schumann et
94 al. 2013; Alfieri et al. 2013) because, in conjunction with the SRTM DEM, it provides a framework
95 within which hydraulic model structures can be assembled. The availability of such datasets reduces
96 significantly the total workload for groups attempting to construct global models, making previously
97 intractable problems manageable for the first time and allowing developers the time to focus on other
98 critical aspects such as efficient numerical schemes and automation.

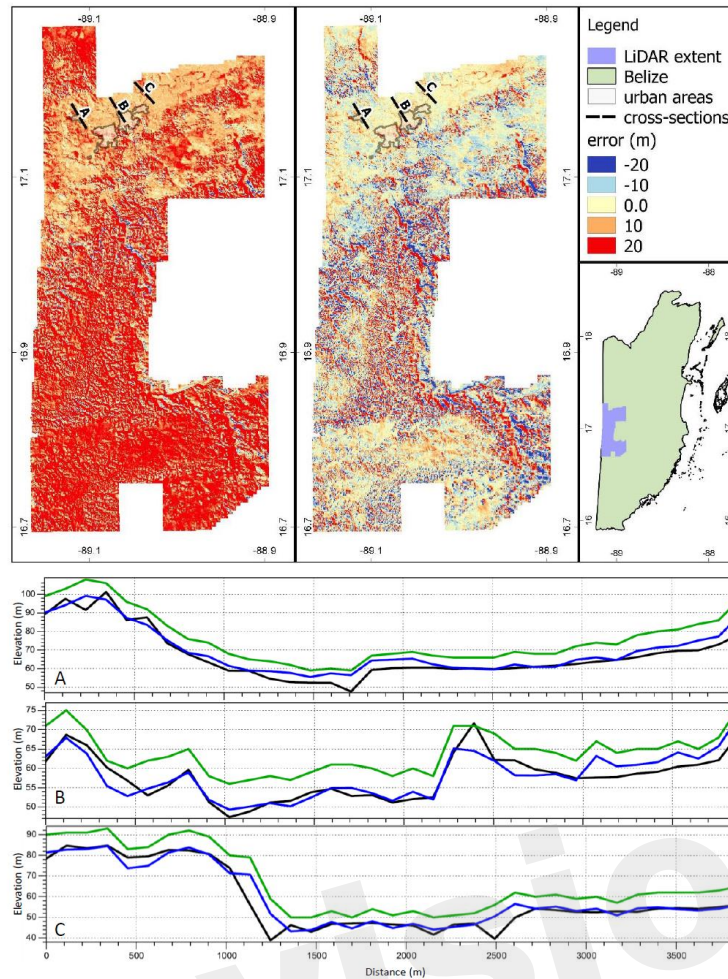
99 However, significant as these achievements may be, the current generation of global DEMs have
100 serious limitations that heavily restrict the skill of models developed around them. Taking SRTM as
101 an example, the critical limitations of the dataset are: a) poor vertical accuracy due to noise or
102 'speckle' (Rodriguez et al. 2006); b) the difficulty in obtaining a bare-earth DTM due to radar
103 reflection from the top of the vegetation canopy; c) the inability to resolve street-scale features in
104 urban areas, resulting in large positive elevation biases in urban areas; d) other systematic errors, such
105 as 'striping', that are a result of the pitch and yaw of the spacecraft during the data collection phase
106 (Rodriguez et al. 2006); and e) the inability of SRTM to resolve the bathymetry of water bodies due to
107 radar reflection from the water surface. These limitations have a highly detrimental effect on both

108 derived hydrography datasets and the simulated flow dynamics of flood hazard models. There is a
109 tendency amongst many users to fixate on the nominal horizontal resolution of DEMs, but for flood
110 modelling it is the vertical accuracy and precision that is critical. This is because the dominant
111 control on the flow of water in a hydraulic model, as in the real world, is the change in elevation of
112 the topography; after all, it is gravity that moves water downslope. The four critical limitations of the
113 SRTM dataset outlined above all concern the vertical accuracy and precision of the DEM, and all can
114 affect simulated flow dynamics adversely. Vertical noise within a DEM will fundamentally affect the
115 propagation of a floodwave because pixels will serve as blockages or sinks. Where noise is random, it
116 can be reduced by resampling the DEM grid to a coarser resolution as the positive and negatively
117 biased pixels cancel when aggregated onto the larger grid. This approach reduces noise, but also
118 reduces the resolution of the DEM and limits its ability to represent small scale features. More
119 challenging still are the elevation biases imparted by vegetation and urban areas. Such biases can be
120 10s of metres and, if left uncorrected, forests and urban areas act as walls or islands that block the
121 flow of water across a floodplain and (erroneously) never flood themselves within the model. As
122 many flood hazard models are used to help assess flood risk, a model that identifies urban areas as
123 always being safe is of little value. Finally, the systematic 'striping' caused by the pitch and yaw of
124 the Space Shuttle itself create false wave like artefacts on the DEM that can corrupt the modelled flow
125 of water across the DEM. It also needs to be noted that SRTM is now quite old (the data were
126 collected in February 2000) and many of the world's floodplains have undergone dramatic change
127 since, mostly because of human development. This is particularly true in developing countries, and
128 there is an increasingly pressing need for a new global topographic mapping mission producing open
129 data.

130 The effect of systematic elevation errors on derived hydrography datasets are equally severe. When a
131 flow direction map is calculated from the DEM, erroneously elevated surfaces caused by areas of
132 vegetation or urbanisation cause errors in the calculated flow directions. This in turn leads to
133 incorrect flow accumulation calculations and stream network locations. The effect can be severe in
134 the case of large forests and cities, leading to grossly misplaced river channels and even missing or

135 invented connections between channels and resultant errors in catchment delineation. The most
136 obvious of these errors can be rectified by painstaking manual editing, as was done for the
137 Hydrosheds dataset (Lehner et al. 2008), but many errors remain that can be hard to identify in a
138 systematic manner. These errors impart structural errors on models that rely upon them for their
139 construction, compounding the DEM-induced errors in flow dynamics discussed above.

140 The errors discussed above have such a marked effect on flood hazard simulations that it has been
141 necessary for practitioners to develop methods that attempt to reduce their severity. One example of
142 this involves attempts to remove vegetation bias from SRTM to produce a bare-earth DTM in forested
143 areas (Baugh et al. 2013). This poses a substantial challenge as the necessary data content is not
144 present in the SRTM data itself, meaning that other datasets are required to quantify the height and
145 location of the vegetation (Simard et al. 2011). Furthermore, because the extent to which the radar
146 pulse penetrates the canopy depends on the density of the vegetation (it is not sufficient to assume the
147 return is always from the top of the canopy), a spatial measure of vegetation density is required.
148 Finally, elevation control points (e.g. ICESat laser altimeter data) are necessary for calibration and
149 validation of the algorithm. Such algorithms can offer significant improvement, as demonstrated in
150 figure 1. However, their effectiveness is limited by the accuracy and precision of the vegetation
151 datasets, which are themselves uncertain, and non-negligible residual errors in the resultant bare-earth
152 DTM are unavoidable; examples of such errors are provided in figure 2 below.



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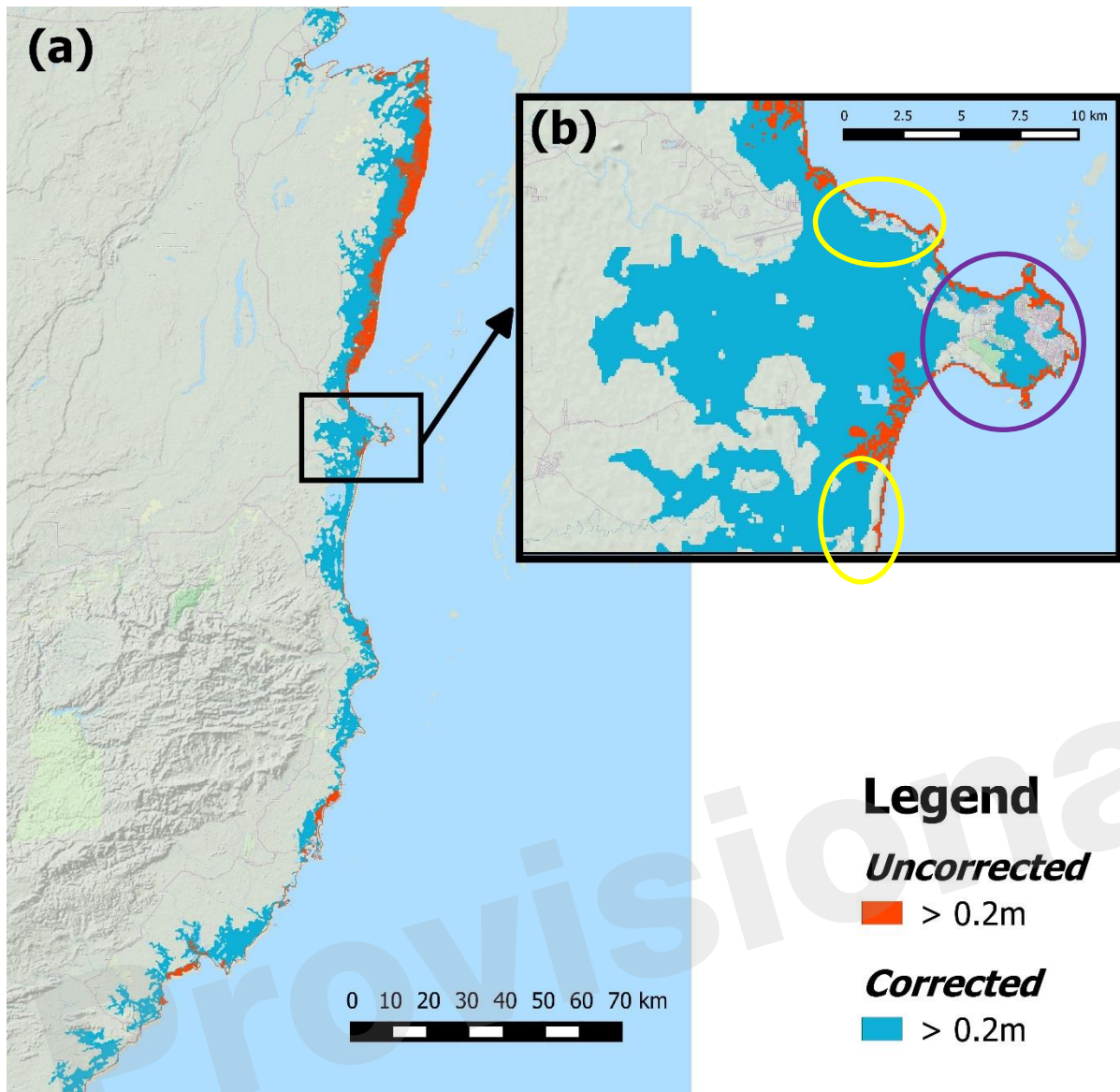
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Figure 1: Comparison of raw SRTM DEM to LIDAR (top left) and corrected SRTM to LIDAR (top centre) for Western Belize. Cross-sections A, B and C transect the Belize River valley and compare the LIDAR elevation profile (black) to the uncorrected SRTM profile (green) and corrected SRTM profile (blue).

Figure 1 shows reduced vertical error following the systematic removal of vegetation bias by comparing corrected and uncorrected SRTM DEMs to a high precision bare-earth DTM produced using 1 m aerial LIDAR data resampled to SRTM resolution. The algorithm employs satellite vegetation height and density datasets (Simard et al. 2011; Schwarz et al. 2004) that estimate vegetation location, height and density to produce an estimated bias layer which is then removed from the SRTM DEM and yields a change in bias from 15.8 to -0.1 m. (Sampson et al. 2015). Figure

165 2 demonstrates the effect of this correction on a simulation of a category 5 storm surge event along
166 the Belize coast. In the uncorrected DEM, the vegetation acts as a virtual ‘sea wall’, preventing the
167 surge waters from penetrating inland to flood areas known to be at risk such as the Belizean coastal
168 mangroves. With the vegetation removed, the coastal wetlands flood, providing a far more plausible
169 realisation of the inundation that one would expect for an event of this magnitude. However, while
170 the improvement is obvious, the transects in figure 1 shows that significant differences still exist
171 between the corrected SRTM DEM and the LIDAR-derived DEM at the local scale due to limitations
172 in the correction method. One key limitation is the resolution of the vegetation datasets (~ 1 km for
173 the vegetation heights and ~250 m for the vegetation density). The yellow circles in figure 2 show
174 areas where the vegetation removal tool failed to resolve and remove ~100m wide strips of mangroves
175 from the SRTM DEM. While the overall removal still allowed water behind the mangrove ‘wall’, this
176 is an example of typical residual vegetation artefacts. It is also known that most of Belize City should
177 be flooded (Belize government engineers and planners, *personal communication*); however dry areas
178 remain due to the residual urban artefacts even after the urban filter is applied to the SRTM DEM
179 (purple circle in figure 2).



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182 Figure 2: Differences in flood extent for a category 5 hurricane storm surge along the Belizean coast using raw vs. vegetation
 183 bias corrected SRTM DEMs

184 There is therefore a clear need for an improved open-access global DEM for global flood hazard
 185 modelling. The value of high resolution terrain data with good vertical precision has long been
 186 recognised at the local scale by the hydraulic modelling community (Fewtrell et al. 2008; Marks &
 187 Bates 2000; Bates et al. 2003; Horritt & Bates 2001; Yu & Lane 2011; Yu & Lane 2006), and the
 188 benefits for global scale models may be even greater. This is because reach scale models often rely
 189 upon manual correction of the DEM using secondary data sources such as surveyed river cross

190 sections; such corrections are not possible on a systematic basis at the global scale because suitable
191 secondary data does not exist for most rivers, and because the scale of the task would render it
192 unfeasible. The DEM is therefore the only source of data used to determine river locations and river
193 bank elevations for most locations within a global model, and it is reasonable to expect any
194 improvements to this dataset to yield substantial improvements to model performance. For example,
195 the representation of flood defences within flood hazard models is known to be critically important
196 (Brandimarte & Di Baldassarre 2012; te Linde et al. 2011; Wesselink et al. 2013), but current large
197 scale models are either forced to assume total failure of defences, or adopt heavily simplified
198 approaches such as masking off urban areas for event scales below a 'defence standard' inferred from
199 socioeconomic data (Feyen et al. 2012) A global DEM of increased horizontal resolution and vertical
200 precision would offer improved representation of micro-topography; if the quality is able to approach
201 that of an aerial LIDAR DEM (> 5 m spatial resolution and 1 m vertical precision), features such as
202 large river levees could be resolved directly. This would lead to the explicit representation of major
203 defence features in large scale models, allowing an improved representation of the flood hazard in
204 protected areas. As even the finest aerial LIDAR DEMs fail to completely capture smaller defence
205 features such as narrow defence walls it is unlikely that any foreseeable global DEM could capture all
206 of the detail necessary for ultra-fine models (Gallien et al. 2014). However, a high quality global
207 DEM could act as a 'base layer' onto which local detail (potentially collected through crowd-sourced
208 platforms such as OpenStreetMap) could be added. Defences are not the only consideration either, as
209 previous studies have shown a step change in model skill for urban areas when the DEM becomes
210 able to resolve individual streets due to correct representation of floodplain connectivity (Fewtrell et
211 al. 2008). A final topic that should be mentioned is cost. According to the Sampson et al. (2015)
212 model, the African 1 in 100 year floodplain covers approximately 7% of the continental area. Scaled
213 to the globe, this gives an approximate 1 in 100 year floodplain area of 35 million km². Assuming
214 some economies of scale, a collection cost of \$200 per km² is plausible and yields a global cost
215 estimate of approximately \$7 billion. As the benefit of the highest resolution data would be most
216 strongly felt in cities, which constitute <0.5% of the Earth's land area (Schneider et al. 2009) but a
217 much larger proportion of the flood risk, one way to significantly reduce the cost of producing such a

218 DEM would be to adopt a hybrid resolution approach where the highest resolution data are collected
219 in urban areas and a lower resolution adopted for rural areas. However, in the context of future
220 annual flood loss estimates that exceed a trillion dollars (Hallegatte et al. 2013), the cost of collecting
221 a high quality global DEM may be justifiable on the basis of its applicability to flood risk modelling
222 alone.

223

224 To conclude, high accuracy and precision DEM data are critical for skilful flood hazard modelling
225 and the limitations with current open access DEM data sets limit significantly our ability to estimate
226 flood inundation and risk for the majority of the planet's surface. There is a clear need (c.f.
227 Schumann et al. 2014) for a concerted global effort to collect or collate a new open access DEM with
228 ~10m resolution and sub-metre scale vertical accuracy for use in a variety of applications. Flood
229 modelling is one such task, but better global DEM data would have wide value for governments,
230 humanitarian organisations, NGOs and industry.

231 **Acknowledgments**

232 We would like to thank John Weishampel, Christine Butterfield and Arlen Chase from the University
233 of Central Florida for supplying the Belize LIDAR data.

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