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RESEARCH ARTICLE

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Key Points:

- Upper ocean oxygenation levels are highly dynamic across OAE 2
- A shallow O₂ oasis in proto-Atlantic is supported by proxy and model
- I/Ca is a proxy for local redox, not for weathering and carbon burial

Supporting Information:

- Table S1 and Text S1

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Citation:

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Upper ocean oxygenation dynamics from I/Ca ratios during the Cenomanian-Turonian OAE 2

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Abstract Global warming lowers the solubility of gases in the ocean and drives an enhanced hydrological cycle with increased nutrient loads delivered to the oceans, leading to increases in organic production, the degradation of which causes a further decrease in dissolved oxygen. In extreme cases in the geological past, this trajectory has led to catastrophic marine oxygen depletion during the so-called oceanic anoxic events (OAEs). How the water column oscillated between generally oxic conditions and local/global anoxia remains a challenging question, exacerbated by a lack of sensitive redox proxies, especially for the suboxic window. To address this problem, we use bulk carbonate I/Ca to reconstruct subtle redox changes in the upper ocean water column at seven sites recording the Cretaceous OAE 2. In general, I/Ca ratios were relatively low preceding and during the OAE interval, indicating deep suboxic or anoxic waters exchanging directly with near-surface waters. However, individual sites display a wide range of initial values and excursions in I/Ca through the OAE interval, reflecting the importance of local controls and suggesting a high spatial variability in redox state. Both I/Ca and an Earth System Model suggest that the northeast proto-Atlantic had notably higher oxygen levels in the upper water column than the rest of the North Atlantic, indicating that anoxia was not global during OAE 2 and that important regional differences in redox conditions existed. A lack of correlation with calcium, lithium, and carbon isotope records suggests that neither enhanced global weathering nor carbon burial was a dominant control on the I/Ca proxy during OAE 2.

1. Introduction

1.1. Marine Environmental Changes During Oceanic Anoxic Events

The concept of oceanic anoxic events (OAEs) was introduced upon the discovery of globally deposited coeval marine organic-rich sediments (black shales) of Cretaceous age, a phenomenon associated directly and indirectly with profound environmental and chemical changes [Schlanger and Jenkyns, 1976; Schlanger et al., 1987; Arthur et al., 1990; Jenkyns, 2003, 2010]. Globally recorded OAEs are recognized as the early Toarcian or T-OAE (~182 Ma) from the Jurassic Period and OAE 1a in the early Aptian (~125 Ma) and OAE 2 (~94 Ma) at the Cenomanian-Turonian boundary from the Cretaceous Period [Ogg and Hinnov, 2012a, 2012b].

During OAE 2, positive carbon isotope excursions (CIEs) in both inorganic and organic carbons are found in different environmental settings [Jarvis et al., 2011]. Because organic carbon has strongly negative $\delta^{13}\text{C}$ values, ranging from -25 to -60 ‰, enhanced rates of organic carbon burial left residual seawater with relatively high carbon isotope values in dissolved inorganic carbon, producing carbonates and organic matter with elevated $\delta^{13}\text{C}$. These positive CIEs, together with macrofossil, microfossil, and nannofossil biostratigraphy, allow correlation of stratigraphic columns from different OAE 2 sections [Tsikos et al., 2004].

Before the CIE started, different proxies suggest that $p\text{CO}_2$ was unusually high [Jarvis et al., 2011], likely introduced by volcanic and hydrothermal activities [Jones and Jenkyns, 2001; Kuroda et al., 2007] related to the formation of the Caribbean and other Large Igneous Provinces [Wignall, 2001; Erba, 2004; Turgeon and Creaser, 2008; Zheng et al., 2013; Du Vivier et al., 2014]. Addition of CO_2 to the atmosphere would have increased global temperature, which enhanced the hydrological cycle and delivered more nutrients to the ocean. In turn, enhanced nutrient availability would have stimulated planktonic productivity, increasing

the oxygen demand in the water column as well as fostering burial of more organic carbon in marine sediments [Weissert, 1989; Jenkyns, 2003; Adams *et al.*, 2010]. A positive excursion of sulfur isotopes in bulk carbonates of Cenomanian-Turonian age indicates that sulfate reduction and enhanced rates of pyrite burial became more significant globally as bottom water euxinia expanded during OAE 2 [Ohkouchi *et al.*, 1999; Owens *et al.*, 2013]. Consistent with this observation, there is considerable biomarker evidence indicating episodic but widespread photic zone sulfide accumulation (euxinia) in the North and South Atlantic regions, the proto-Indian Ocean, and in the Tethys [Sinninghe Damsté and Koster, 1998; Kuypers *et al.*, 2002; Pancost *et al.*, 2004; Forster *et al.*, 2008; van Bentum *et al.*, 2009].

Beyond the sulfur, carbonate, and organic matter isotope records, the inorganic elemental proxies applied to OAE 2 are mostly restricted to concentrations and elemental ratios of redox-sensitive metals in shales (e.g., Mn, Mo, V, Fe, U, Co, Ce, and Cd), although metal isotopes are being increasingly used. Trace metals precipitate as sulfides under anoxic and euxinic conditions, some with valence changes (e.g., Mo and V) and some without, such as Zn [Hetzel *et al.*, 2009]. During OAE 2, some trace metals, such as As, Bi, Cd, Co, Cr, Cu, Mo, Ni, Sb, Tl, V, and U, were concentrated in organic matter or precipitated with sulfide, suggesting euxinia in the sediments and/or in the water column [Arthur *et al.*, 1990; Kuypers *et al.*, 2002; Jenkyns, 2010], as documented for the proto-Atlantic [Kuypers *et al.*, 2002; Kolonic *et al.*, 2005; Brumsack, 2006; Forster *et al.*, 2008; Hetzel *et al.*, 2009; Tribouillard *et al.*, 2012; van Helmond *et al.*, 2014; Little *et al.*, 2015] and for southern Europe [Scopelliti *et al.*, 2006, 2008; Turgeon and Brumsack, 2006]. Molybdenum isotope data [Westermann *et al.*, 2014] and uranium isotope data [Montoya-Pino *et al.*, 2010] also indicate expanded bottom water euxinia during the OAE 2. Although these geochemical species are interpreted as paleoredox tracers, they may also have responded to changes in hydrothermal flux that may have further contributed to metal enrichments during OAE 2 [Orth *et al.*, 1993; Snow *et al.*, 2005; Elrick *et al.*, 2009; Jenkyns, 2010; Eldrett *et al.*, 2014].

The details of how the global ocean evolved from relatively widespread oxidizing conditions, with only localized anoxia in restricted marine settings and areas of particularly high primary production, to globally extensive anoxia [Pancost *et al.*, 2004; Hetzel *et al.*, 2009; Montoya-Pino *et al.*, 2010; Owens *et al.*, 2013] remain unclear for OAE 2 and OAEs more generally. Newly developed geochemical techniques utilizing the biophilic element iodine in carbonates help to bridge this gap and shed novel light on the geographic patterns and controls on the development of anoxia.

1.2. I/Ca as a Paleoredox Proxy

Although iodine can exist in several oxidation states, iodide (I^-) and iodate (IO_3^-) are the thermodynamically stable inorganic forms in seawater. The standard reduction potential of IO_3^-/I^- is very close to that of O_2/H_2O [Rue *et al.*, 1997; Harris, 2006], making iodine one of the first elements to respond to oceanic deoxygenation. Most of the iodine in the modern well-oxygenated ocean occurs as iodate [Wong, 1995; Farrenkopf *et al.*, 1997; Campos *et al.*, 1999; Waite *et al.*, 2006]. Because only iodate precipitates with carbonate, the simple presence of carbonate-associated iodine requires locally oxic conditions in the water column. Iodate is, however, almost completely reduced to iodide in all investigated modern anoxic basins and oxygen minimum zones where dissolved O_2 is less than $3 \mu M$ [Wong and Brewer, 1977; Wong *et al.*, 1985; Luther and Campbell, 1991; Farrenkopf *et al.*, 1997; Rue *et al.*, 1997; Farrenkopf and Luther, 2002]. Consequently, a drop in iodate concentration to ~ 0 indicates strong deoxygenation within the local water column.

I/Ca ratios in calcite increase linearly with iodate concentrations in the precipitating medium, but iodide does not incorporate into carbonate [Lu *et al.*, 2010]. The mechanism of iodate incorporation is unclear but may involve substitution for the carbonate ion and/or the presence of lattice defects, perhaps analogous to the mechanism proposed for sulfate incorporation into carbonate [Staudt and Schoonen, 1995]. The lack of iodide incorporation is likely due to the large ionic radius of iodide relative to iodate.

Most pelagic carbonate is produced within the upper levels of the water column as coccolith and planktonic foraminiferal calcite, which in recent environments falls to the seafloor roughly in the proportion of 1:1 [Broecker and Clark, 2009]. Hence, a bulk carbonate chemical signal should reflect conditions in near-surface waters. Although the percentage of biogenic carbonate produced at the seafloor is generally very small, the influence of benthic calcifiers cannot be ruled out. For these reasons, the bulk carbonate I/Ca ratio in this work is taken as an indicator for the integrated conditions in the upper ocean water column.

Once oxic conditions are established, iodine speciation at a given location is secondarily influenced by regional mixing. The kinetics of iodide oxidation are not well constrained [Luther *et al.*, 1995] but are currently estimated to be slow, with the lifetime of iodide in oxygenated waters ranging from months to several decades [Chance *et al.*, 2014]. Since iodide is not oxidized instantaneously, upwelling of anoxic (iodide-rich) waters from anoxic basins or oxygen minimum zones (OMZs) to well-oxygenated surface waters results in lower iodate concentrations compared to those in the open ocean, promoting accumulation of bulk carbonate with relatively low I/Ca ratios. Such a decrease in bulk I/Ca does not necessarily represent lower dissolved oxygen levels in the mixed layer but instead reflects redox conditions in the underlying water masses that are in direct exchange with near-surface waters.

From a global perspective, local I/Ca ratios are also expected to reflect potential significant shifts in the size of the marine iodine reservoir. The contemporary marine iodine budget is not accurately constrained. However, the iodine fluxes (e.g., river and mantle inputs, output of organic matter burial) in and out of seawater are estimated to be 1 to 2 orders of magnitude lower than the amount of iodine associated with biological production [Lu *et al.*, 2010, and references therein]. The biological pump results in some iodate loss from mixed-layer waters in high-productivity regions. Such surface depletions have been observed at stations off Hawaii, Bermuda [Campos *et al.*, 1996], the Weddell Sea [Bluhm *et al.*, 2011; Campos *et al.*, 1999], Mediterranean Sea [Tian *et al.*, 1996], Arabian Sea [Farrenkopf and Luther, 2002], and an Antarctic coastal site [Chance *et al.*, 2010]. Most of this iodine is released back into the seawater during the decomposition of organic matter in an oxic water column [Lu *et al.*, 2010, and references therein], but organic carbon burial results in a large flux of iodine to the sediments. Consequently, enhanced primary production should export more organic iodine out of the mixed layer, lowering total iodine concentration and hence lowering I/Ca in bulk carbonate. If the burial of organic carbon can lead to global iodine drawdown, it is anticipated that I/Ca decreases will generally correlate with increases in $\delta^{13}\text{C}$.

Bulk carbonate I/Ca signals may be affected by other factors in addition to the local upper ocean redox conditions. Mixing with diagenetic carbonate or recrystallization of primary carbonate in anoxic pore waters should lower the I/Ca values, since iodide, the only inorganic iodine species in anoxic marine or pore waters [Wong and Brewer, 1977; Wong *et al.*, 1985; Kennedy and Elderfield, 1987a, 1987b; Luther and Campbell, 1991; Farrenkopf *et al.*, 1997; Rue *et al.*, 1997; Farrenkopf and Luther, 2002], cannot be incorporated into carbonate [Lu *et al.*, 2010]. Therefore, carbonates with obvious diagenetic features should be avoided for measurements of I/Ca.

1.3. Multiproxy Comparisons

New I/Ca data presented and discussed in this study represent several end-member locations recording OAE 2, covering different paleowater depths, paleolatitudes, organic carbon contents, and accumulation rates (Figure 1 and Table 1). Sulfur isotope data [Adams *et al.*, 2010; Owens *et al.*, 2013] are compared with our I/Ca records to further constrain the global redox conditions. Relatively low I/Ca values found at multiple sections recording the OAE may indicate widespread oxygen-depleted conditions. For example, during the early Toarcian OAE, a section of shallow-water carbonates recorded an I/Ca decrease of $\sim 4 \mu\text{mol/mol}$ in phase with the onset of positive carbon isotope excursion (CIE) and coincided tightly with the beginning of a large shift in the sulfur isotope composition of carbonate-associated sulfate [Lu *et al.*, 2010; Gill *et al.*, 2011].

The $\delta^{34}\text{S}_{\text{CAS}}$ can provide useful information about changes in paleoredox, but it could be affected by multiple environmental changes, such as changes in the input from continental weathering, volcanism, and hydrothermal activity; pyrite and evaporite deposition; and the availability of sulfate, iron, and organic matter [e.g., Paytan *et al.*, 2004]. However, recording a positive $\delta^{34}\text{S}_{\text{CAS}}$ excursion in multiple ocean basins over a short event is nearly impossible without the burial of isotopically light sulfur, most plausibly from a global increase in pyrite burial. Importantly, although I/Ca and $\delta^{34}\text{S}$ are expected to show general similarities in their behavior during major oxygenation changes, they will not necessarily covary precisely, due to their distinct biogeochemical behaviors [Hardisty *et al.*, 2014]. In the modern ocean, the residence time for sulfate is 10–15 Myr [Walker, 1986], which is much longer than the ~ 300 kyr estimated for iodine [Broecker and Peng, 1982], although this residence time is likely to have been different during the OAE 2 since oceanic sulfate concentrations perhaps were $\sim 1/4$ of the modern value (7 mM versus 28 mM) [Paytan *et al.*, 2004; Owens *et al.*, 2013]. Generally, however, $\delta^{34}\text{S}_{\text{CAS}}$ should illustrate dominantly a global signal and I/Ca a local one.

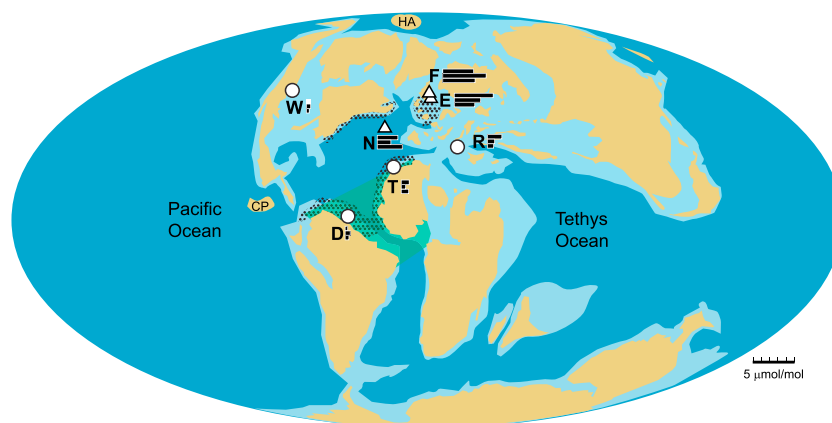


Figure 1. Global paleogeographic map of OAE 2 sections. The yellow represents the continents, the light blue stands for shallow oceans, and the dark blue for deep oceans. The black dots mark the modeled upwelling regions from *Topper et al.* [2011], and the green shading shows areas of euxinic conditions in the water column given in *Jenkyns* [2010]. The studied sites are grouped into oxic (white triangles) and hypoxic (white circles) sites, and the site names are abbreviated. See Table 1 for abbreviations for site names. Average I/Ca values during pre-CIE, CIE, and post-CIE intervals are shown as uniformly scaled bars from bottom to top for individual sites. CP is Caribbean Plateau, and HA is High Arctic Large Igneous Province.

Calcium and lithium isotopes are important markers for crustal weathering rates during OAE 2 [*Blättler et al.*, 2011; *Pogge von Strandmann et al.*, 2013]. These isotopic records are also examined to illuminate any potential influence of weathering on the marine iodine budget and I/Ca signals. Finally, the oceanographic I/Ca pattern is linked to the oxygenation conditions modeled in the Grid Enabled Integrated Earth System Model (GENIE) [*Monteiro et al.*, 2012].

2. Study Sites

New I/Ca data are reported from seven sections and compared with the previously published data from Eastbourne, UK (Table 1) [*Lu et al.*, 2010; *Tsikos et al.*, 2004]. The sections are mostly located either within or peripheral to the proto-Atlantic Ocean (Figure 1). This sample set covers a variety of paleowater depths and lithologies from shallow-water platform carbonates to pelagic chalks and shales. Material from Raia del Pedale, Italy; Eastbourne, UK; South Ferriby, UK; and Newfoundland Drifts, northwestern Atlantic are mostly organic-lean carbonates, although the latter two sites contain thin (centimeter-scale) black organic-rich shales. The Aristocrat Angus core in the U.S. Western Interior Seaway (WIS) and at Demerara Rise, offshore Suriname and the Tarfaya (Morocco) site, have relatively elevated total organic carbon (TOC) in marls and shales (commonly >5%) throughout the studied interval.

Table 1. OAE 2 Sections Used for I/Ca Measurements and Multiproxy Comparison^a

Label	Location	Deposition Environment	Lithology	Reference
R	Raia del Pedale, Italy	Shallow water	Limestone	<i>Parente et al.</i> [2008]
W	Aristocrat Angus core, Western Interior Seaway	Shallow water	Limestone/shale	<i>Joo and Sageman</i> [2014]
D	Demerara Rise, ODP 1258	Low-latitude, Pelagic	Organic-rich limestone/Black shale	<i>Hetzel et al.</i> [2009]; <i>Erbacher et al.</i> [2005]
T	Tarfaya, Morocco	Low-latitude, Pelagic	Organic-rich Chalk/Black shale	<i>Jenkyns et al.</i> [2007]
E	Eastbourne, UK	Midlatitude, Pelagic	Chalk	<i>Pearce et al.</i> [2009]
F	South Ferriby, UK	Midlatitude, Pelagic	Chalk/Black shale	<i>Jenkyns et al.</i> [2007]
N	Newfoundland Drifts, IODP U1407	Midlatitude, Pelagic	Chalk/Black shale	<i>Expedition 342 Scientists</i> [2012]

^aAbbreviations are used in the site map.

3. Analytical Methods

3.1. Sample Preparation and Measurements

Powdered samples (3–5 mg) from different lithological settings were weighed out on a microgram balance and then rinsed with deionized water to remove dissolvable iodine salt potentially attached to the surface of the samples. Next, a volume of nitric acid (3%) was calculated according to sample weight, and the acid was added to each sample in a sonicating bath to allow complete dissolution of the carbonate. Residual noncarbonate impurities were removed by centrifuging. The final solutions for the inductively coupled plasma–mass spectrometry (ICP-MS) measurements contain approximately 50 ppm Ca and 0.5% tertiary amine to help stabilize iodate [Schnetger and Muramatsu, 1996] and internal standards. For each batch of measurements, pure potassium iodate was dissolved and diluted as a calibration standard. I/Ca was measured on a quadrupole ICP-MS (Bruker M90) at Syracuse University. The sensitivity of I-127 is tuned to about 80–100 kcps for a 1 ppb standard. The precision of ^{127}I is typically better than 1% and is not reported individually for each sample. The long-term accuracy is guaranteed by repeated measurements of the reference material JCP-1 [Lu et al., 2010]. The detection limit of I/Ca is usually below 0.1 $\mu\text{mol/mol}$. I/Ca values generated in this process are specific to carbonate, with minor influence from clay, silicate, and organic matter. The liberation of iodine attached to noncarbonate phases during chemical analysis can artificially increase the I/Ca values. This process is particularly important when dealing with sediments relatively high in TOC, because marine organic matter is enriched in iodine.

3.2. Earth System Modeling

As an aid to the interpretation of our I/Ca records, we make use of a pair of simulations carried out in a 3-D ocean circulation-based Earth System Model (“cGENIE”) [Ridgwell et al., 2007]. For this paper, the climatology and continental arrangements are configured for the Cenomanian-Turonian as described in Monteiro et al. [2012]. Carbon and other biogeochemical cycles in the ocean include PO_4^{3-} , NO_3^- , and NH_4^+ as nutrients controlling biological export production, plus O_2 , NO_3^- , and SO_4^{2-} as potential oxidants [Ridgwell et al., 2007; Monteiro et al., 2012]. We ran two experiments: one representing potential pre-OAE 2 redox conditions and assuming a modern ocean PO_4^{3-} inventory and $\times 2$ preindustrial CO_2 in the atmosphere and another assuming syn-OAE 2 conditions, with $\times 2$ PO_4^{3-} and $\times 4$ CO_2 . These combinations of oceanic PO_4^{3-} and atmospheric CO_2 determined by Monteiro et al. [2012] yield the best possible fit to available indicators of ocean floor anoxia and photic zone euxinia. We ran both experiments for 20 kyr to achieve full steady state with respect to patterns of ocean oxygenation and nitrogen cycling. The model code and experimental configurations are identical to those described in Monteiro et al. [2012] and can be replicated via the instructions given in the supporting information.

4. Results and Discussions

4.1. Influence of Diagenesis and Organically Bound Iodine

Bulk carbonate I/Ca signals may be affected by secondary factors in addition to the local upper ocean redox conditions. Mixing with diagenetic carbonate or recrystallization of primary carbonate in anoxic pore waters should lower the I/Ca values, since iodide in anoxic waters cannot be incorporated in carbonate. It is possible that all of the I/Ca values are influenced by diagenesis to some extent. However, the consistency demonstrated by multiproxy and data-model comparisons across different depositional settings suggests that bulk carbonate I/Ca can still serve as a reasonably reliable redox proxy for the suboxic window. Site 1258 on Demerara Rise is the only locality where diagenetic carbonate is common [Erbacher et al., 2004]. The CIE is not very well expressed in the carbonates at the Tarfaya section, indicating possible diagenesis, but the carbonate record in Eastbourne appears largely unaffected by diagenesis and has proved an invaluable geochemical archive [Tsikos et al., 2004]. Minor dolomite is present in the basal part of the Raia del Pedale section but below the CIE interval. As demonstrated below, Tarfaya and the Western Interior Seaway (WIS) sites both have relatively high TOC contents, but the corresponding I/Ca values are much lower than those at sites dominated by low-TOC chalk, suggesting that organic iodine does not always significantly affect I/Ca. Furthermore, the lowest I/Ca values at Tarfaya and in the WIS correlate with high TOC during the OAE. These observations strongly argue against organic matter being an important contaminant of the carbonate I/Ca signal, particularly since our method employs brief exposure to only dilute acid.

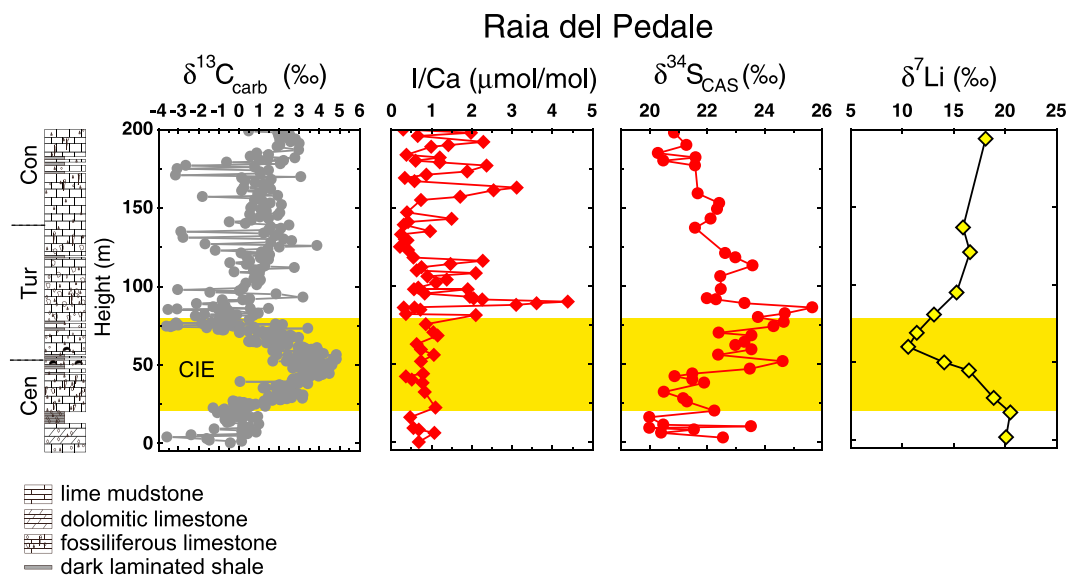


Figure 2. Chemostratigraphic plot for OAE 2 at Raia del Pedale. I/Ca data are from this study, $\delta^{13}\text{C}$ and $\delta^{34}\text{S}_{\text{CAS}}$ data are from Owens *et al.* [2013], and $\delta^7\text{Li}$ are from Pogge von Strandmann *et al.* [2013]. The stratigraphic column on the left shows the facies variations [Owens *et al.*, 2013], and the yellow boxes highlight the positive CIE.

4.2. Shallow-Water Sites

At all seven sections, the OAE is defined by the positive CIE and ensuing recovery. Raia del Pedale, Italy, is an organic-lean, carbonate-platform section comprising gray lime mudstones containing benthic foraminifera and thick-shelled fossils such as rudists, likely deposited in waters only a few meters deep [Parente *et al.*, 2008] (Figure 1). Although the pre-OAE interval is not exposed in outcrop, the OAE and post-OAE intervals were sampled in detail (Figure 2). During the pre-CIE and CIE, I/Ca values are consistently low at $\sim 1 \mu\text{mol/mol}$, indicating that the local near-surface waters had relatively low iodate concentrations and were likely proximal to and were mixing with low-oxygen iodide-rich waters brought onto the platform. These conditions were stable locally before and during OAE 2, as indicated by the small variability in I/Ca. Immediately above the level of the CIE, I/Ca briefly increases to above $4 \mu\text{mol/mol}$ and then returns to low values similar to those of the OAE, but with scattering between ~ 0.5 and $2.5 \mu\text{mol/mol}$. I/Ca values average at $1.2 \mu\text{mol/mol}$, which are lower than those at pelagic shelf sea sites such as Eastbourne [Lu *et al.*, 2010].

The $\delta^{34}\text{S}_{\text{CAS}}$ plot for Raia del Pedale shows a gradual increase of up to 6‰ during and slightly post-CIE and then declines further up section. The timing of the reversal of $\delta^{34}\text{S}_{\text{CAS}}$ trend coincides precisely with the rapid recovery in I/Ca at the end of the CIE (Figure 2), suggesting that a more oxygenated water mass developed post-OAE at this site, conforming to global trends of lower pyrite burial rates. There is no notable relationship between I/Ca ratios and Li isotope profiles [Pogge von Strandmann *et al.*, 2013], suggesting that weathering rates do not control the iodine trend at this locality.

The Aristocrat Angus core in Western Interior Seaway (WIS) records the well-established positive global carbon isotope excursion during the OAE 2 [Sageman *et al.*, 2006; Joo and Sageman, 2014]. Sulfur isotope records have been generated from the Portland core [Adams *et al.*, 2010]; wt % TOC does not covary systematically with these isotopic records [Sageman *et al.*, 1998; Meyers *et al.*, 2005]. The WIS represents a unique depositional environment among the studied sites. It was a relatively restricted shallow-water seaway (100 to 200 m water depth [Kauffman and Caldwell, 1993]), where local iodine concentration and speciation should have been very sensitive to organic matter burial and redox conditions. I/Ca ratios from the Aristocrat Angus core [Joo and Sageman, 2014], WIS, are all relatively low ($< 1.2 \mu\text{mol/mol}$; Figure 3). The I/Ca ratios peak at the early stage of the CIE at levels marked by relatively low TOC content. This observation may indicate a short period characterized by better oxygenated local waters, known as the Plenus Cold Event/Benthic Oxidic Event. This event has been linked to climatic cooling and the presence of more oxygenated waters in other regions during the same time interval and was characterized by a

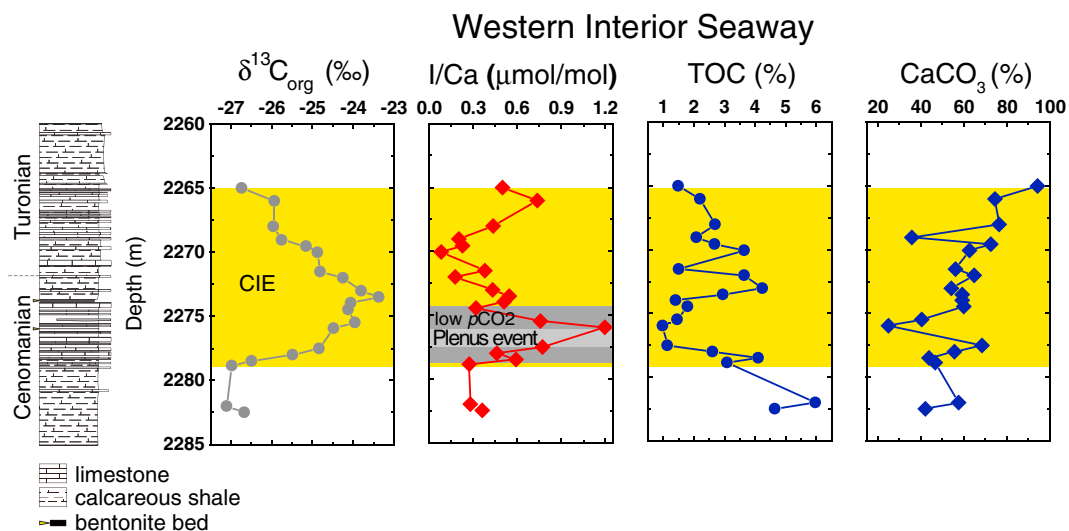


Figure 3. Preliminary I/Ca data from the Aristocrat Angus core in the Western Interior Seaway, plotted with wt % TOC, wt % CaCO₃, and δ¹³C_{org} [Joo and Sageman, 2014]. The stratigraphic column is revised from Joo and Sageman [2014]. The yellow boxes mark the CIE.

temporary influx of benthic foraminifera in successions otherwise generally barren of benthic fauna [Gale and Christensen, 1996; Keller and Pardo, 2004; Forster et al., 2007; Jarvis et al., 2011; Alegret and Thomas, 2013; Eldrett et al., 2014]. Lower I/Ca and higher TOC following the Plenus Cold Event suggest a return to local more poorly oxygenated conditions during the latter part of the global OAE.

4.3. Low-Latitude Pelagic Sites

The tropical proto-Atlantic, particularly the northern South Atlantic, is thought to have hosted the most reducing water masses during OAE 2, based on the relative abundance of biomarkers for phototrophic sulfide oxidizers [Sinninghe Damsté and Koster, 1998; Kuypers et al., 2002]. Our study centered on two locations: Demerara Rise and Tarfaya, Morocco, which are characterized by high TOC levels and serve as tropical peri-equatorial end-member sites (Figure 1). Both sites are within the modeled upwelling region [Topper et al., 2011] dominated by deepwater euxinia and episodic photic-zone euxinia [Kuypers et al., 2002; van Bentum et al., 2009]. Shallow euxinic conditions or upwelling of euxinic deep water should have limited the iodate concentrations in surface waters significantly and set up a steep gradient of iodate concentration between the sea surface and the base of the photic zone.

The I/Ca ratios at Tarfaya are between 0.14 and 1.6 µmol/mol (Figure 4), which are significantly lower than the range at Eastbourne (~2–5 µmol/mol). The I/Ca profile generally changes in the opposite direction to that of carbon isotopes at the same time. During the peak CIE, I/Ca ratios are relatively low, and δ¹⁵N values illustrate a broad positive excursion, which may indicate regional upwelling of anoxic waters that promoted both iodate reduction and denitrification and/or anammox processes [Jenkyns et al., 2007].

Site 1258 on Demerara Rise is currently located at a water depth of over 3 km below modern sea level, although the exact depth of this submarine feature during the mid-Cretaceous remains unresolved. The majority of I/Ca ratios at Site 1258 are below 0.6 µmol/mol, with the lowest average ratio among all the sections (Figure 1). The stratigraphic trend for I/Ca is broadly the inverse of the trend for carbonate content (Figure 5), which is likely a result of authigenic carbonate diluting/dominating the bulk rock I/Ca signal. Sediments with lighter colors in the studied interval reflect higher carbonate content, identified as layers of significant diagenetic calcite growth [Erbacher et al., 2004]. The high CaCO₃% and near-zero I/Ca values at core depths of 421 and 428 m also coincide with peaks of bulk sediment Mn concentration, interpreted to reflect formation of authigenic Mn-rich carbonate [Hetzel et al., 2009]. Diagenetic carbonate that precipitated in anoxic pore waters does not incorporate iodine, because iodide is the only iodine species in reducing pore water [Kennedy and Elderfield, 1987a, 1987b], and as discussed above, the calcite

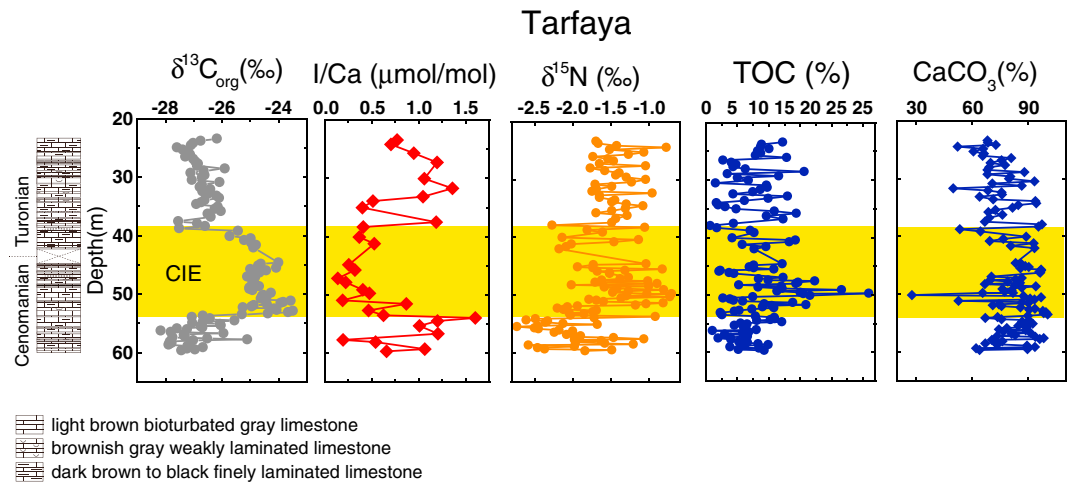


Figure 4. Geochemical and stable isotope data for OAE 2 at Tarfaya (core S57). $\delta^{13}\text{C}_{\text{org}}$ marks the CIE well and covers short periods before and after the CIE. The I/Ca data from bulk carbonate are generated in this study, compared with the $\delta^{15}\text{N}$, wt % TOC, and wt % CaCO_3 [Jenkyns et al., 2007]. The highest TOC (%) correlates with the CaCO_3 (%) minimum and $\delta^{15}\text{N}$ maximum during the CIE. The stratigraphic column shows variable limestone, revised from Jenkyns et al. [2007]. The CIE is bracketed in yellow boxes.

structure does not accommodate this ion [Lu et al., 2010]. Hence, a significant quantity of diagenetic calcite has likely compromised the primary I/Ca record of this particular locality.

However, other proxies suggest that euxinic water masses were present at this site [Hetzel et al., 2009], which may indicate that low I/Ca ratios, if not solely due to diagenesis, also represent anoxic conditions near the surface, either by reduction in the local area or by mixing with underlying or laterally adjacent anoxic water masses. Euxinic conditions, below the surface waters, are suggested at Site 1258 by the Fe, S, and Mo data [Hetzel et al., 2009], and $\delta^{15}\text{N}$ values suggest upwelling of reduced nitrogen species that fueled planktonic productivity [Higgins et al., 2012]. Consistent with these indicators, the interval analyzed in our study consists of finely laminated black shales with locally developed phosphatic limestone nodules [Erbacher et al., 2005].

4.4. Midlatitude Pelagic Sites

New I/Ca data have been generated from a northeast European shelf sea pelagic sequence (South Ferriby in the UK) and from a lithologically similar sequence in the Newfoundland Drifts in the northwest Atlantic [Expedition 342 Scientists, 2012]. The South Ferriby section contains mostly organic-lean chalk, except for an

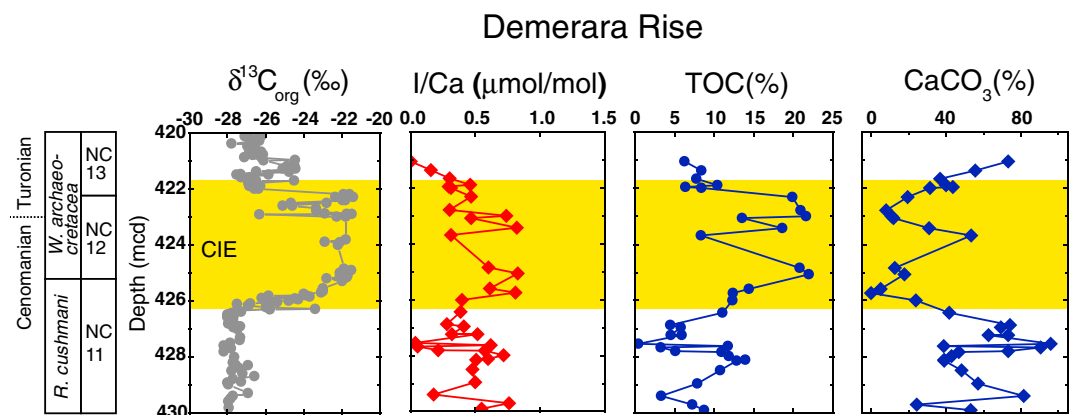


Figure 5. I/Ca for ODP Site 1258 on Demerara Rise, compared with $\delta^{13}\text{C}_{\text{org}}$, wt % TOC [Erbacher et al., 2005], and wt % CaCO_3 (iodp.tamu.edu). The biostratigraphic column to the left is from Erbacher et al. [2005]. The scale is meter composite depth.

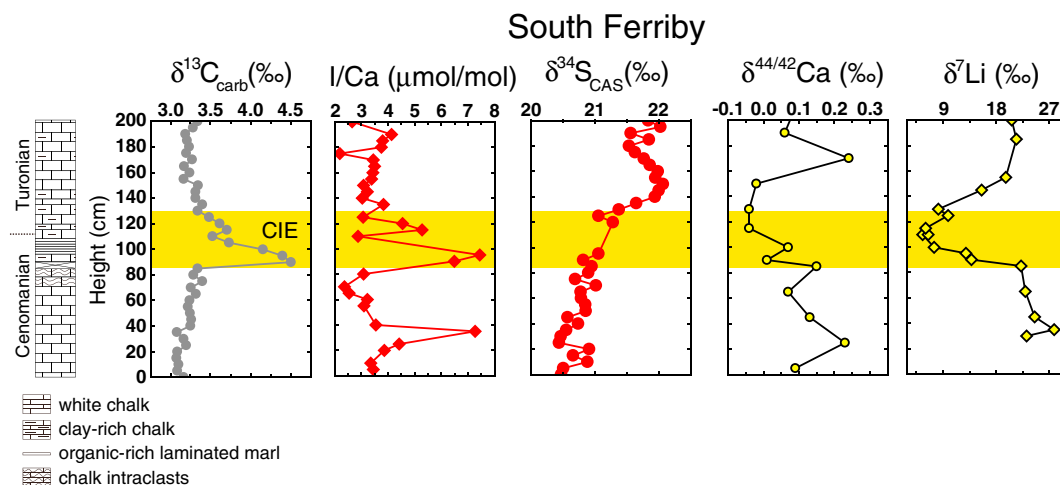


Figure 6. I/Ca data from South Ferriby, compared with $\delta^{13}\text{C}_{\text{carb}}$, $\delta^{34}\text{S}_{\text{CAS}}$, Ca, and Li isotopes [Blättler *et al.*, 2011; Owens *et al.*, 2013; Pogge von Strandmann *et al.*, 2013]. The stratigraphic column is adapted from Jenkyns *et al.* [2007]. The yellow boxes bracket the CIE. A hiatus appears at the beginning of the CIE at ~85–90 cm and is highlighted in the stratigraphic column as a crossed box.

interval of ~10 cm thick black shale, where TOC rises to 8% [Jenkyns *et al.*, 2007]. The onset and main part of the CIE are missing at South Ferriby due to a hiatus (Figure 6). The baseline of I/Ca ratios shows a slight decrease starting below the CIE and a gradual increase above the CIE. The preevent baseline varies between 2 and ~4 $\mu\text{mol/mol}$, which is very similar to the range at Eastbourne, indicating an oxygenated water mass. Notably, much higher I/Ca spikes are observed immediately below three clay-rich layers in this section (Figure 6). The spikes became increasingly muted during the recovery from the peak CIE. The highest I/Ca value (7.4 $\mu\text{mol/mol}$) found here is the highest among all OAE 2 carbonates. These spikes may be related to local invasion of well-oxygenated water masses with high iodate concentrations.

The $\delta^{34}\text{S}_{\text{CAS}}$ increases during the CIE and then starts to decrease at ~140 cm, postdating the CIE, where the I/Ca spikes disappear and the baseline rises. Both Ca and Li isotopes indicate enhanced global weathering [Blättler *et al.*, 2011; Pogge von Strandmann *et al.*, 2013], which may have increased river input of iodine into the ocean and generated higher I/Ca values. These negative Ca and Li isotope trends, however, broadly correspond to the interval of low I/Ca values in most sections, with the exception of a few spikes at South Ferriby. These oscillations in I/Ca are more likely the result of local episodic oceanographic changes rather than changes in the global iodine budget. Unlike the WIS and Eastbourne sites, the Plenus Cold Event is not recorded from the South Ferriby site. The South Ferriby section is, however, condensed by a factor of ~20 compared to Eastbourne [Gale *et al.*, 1993], and a hiatus is present near the level of onset of the CIE in this section, making it likely that the Plenus Cold Event was lost from the sedimentary sequence [e.g., Hart and Leary, 1989].

The lithostratigraphy at Site U1407, Newfoundland Drifts, comprises a thick sequence of nannofossil chalk, varying from white to dark gray in color, above and below the black shale (~0.4 m near the depth of 232 m). A hiatus appears within the CIE interval at ~231–232 m. Unpublished carbon isotope data (C. Junium) indicate that the CIE bracketing the black shale outlines the OAE 2 interval between 229 and 233 m (Figure 7). I/Ca ratios mostly fluctuate at 2–4 $\mu\text{mol/mol}$ before the CIE, with lower values occasionally observed coinciding with dark chalk. The I/Ca ratios decrease right at the onset level of the CIE, coincident with the base of the black shale. The recovery of I/Ca starts with values below 1 $\mu\text{mol/mol}$ and rises slowly to pre-OAE background levels near 3 $\mu\text{mol/mol}$. These data suggest that at this site, the proximal water masses that mix with the local surface waters were generally oxygenated before the OAE, whereas oxygen was rapidly removed from the local surface waters, where the iodine signal is captured, at the onset of the event itself, before slowly returning to background levels.

Eastbourne is among the most studied OAE 2 sites and was included in the first I/Ca investigation of the ancient ocean [Lu *et al.*, 2010]. However, several important observations have since emerged from

Newfoundland Drifts

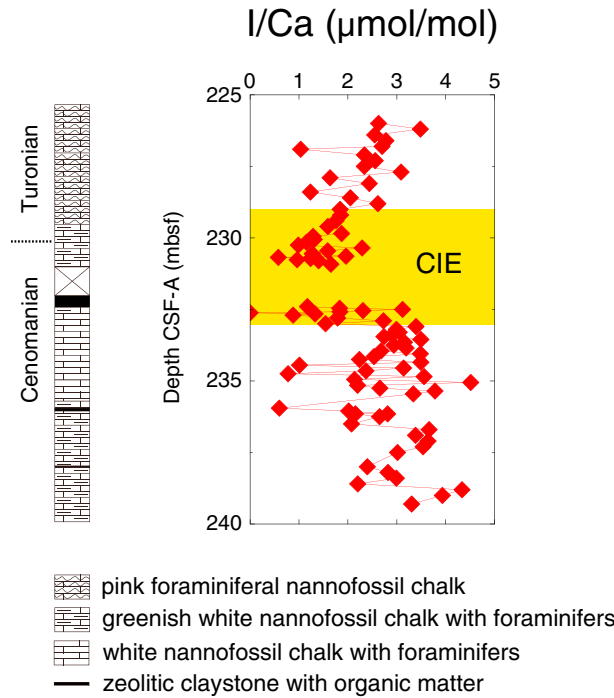


Figure 7. I/Ca record of Site U1407, IODP X342 on Newfoundland Drifts. The stratigraphic log is adapted from *Expedition 342 Scientists* [2012]. A hiatus at ~231–232 is represented by a crossed box. The mbsf = meters below seafloor.

found at Eastbourne at a level close to the suggested base of the Plenus interval indicate that the spatial extent of oxygenated waters may have been more widespread than previously assumed during this event, although upper ocean geochemical conditions may not have evolved synchronously with bottom water conditions.

Eastbourne

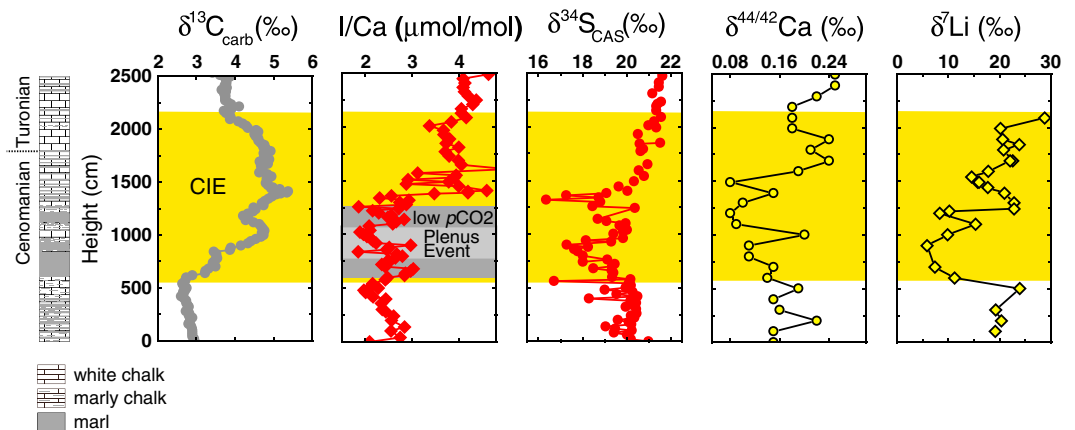


Figure 8. Geochemical and isotope data for Eastbourne. The $\delta^{13}\text{C}_{\text{carb}}$ marks the CIE well, and I/Ca and $\delta^{34}\text{S}_{\text{CAS}}$ indicate redox changes during the studied interval. Plenus Cold Event, also known as the Benthic Oxidic Event, is bracketed in light gray box [Tsikos et al., 2004; Owens et al., 2013]. The ranges of low $p\text{CO}_2$ and Plenus Cold Event are adapted from Jarvis et al. [2011]. The $\delta^{44/42}\text{Ca}$ and $\delta^7\text{Li}$ are regarded as proxies of continental weathering rates [Blättler et al., 2011; Pogge von Strandmann et al., 2013].

multiproxy comparisons that have refined our view of those data [Blättler et al., 2011; Owens et al., 2013; Pogge von Strandmann et al., 2013]. Relatively high I/Ca values and an excursion to lower $\delta^{34}\text{S}_{\text{CAS}}$ values are reported from Eastbourne at levels corresponding with the lower part of the CIE (Figure 8) [Lu et al., 2010; Owens et al., 2013]. These geochemical data likely record reoxygenation related to the Plenus Cold Event, which occurred during the early phase of OAE 2 and was initially proposed for the European Chalk Sea and the Western Interior Seaway based on southward invasion of boreal faunas [Gale and Christensen, 1996; Keller and Pardo, 2004; Forster et al., 2007]. However, $p\text{CO}_2$ records based on stomata and $\Delta^{13}\text{C}$ indicate that the cooling and CO_2 drawdown were likely global during the first third of the CIE [Friedrich et al., 2006; Barclay et al., 2010; Jarvis et al., 2011]. The high I/Ca values

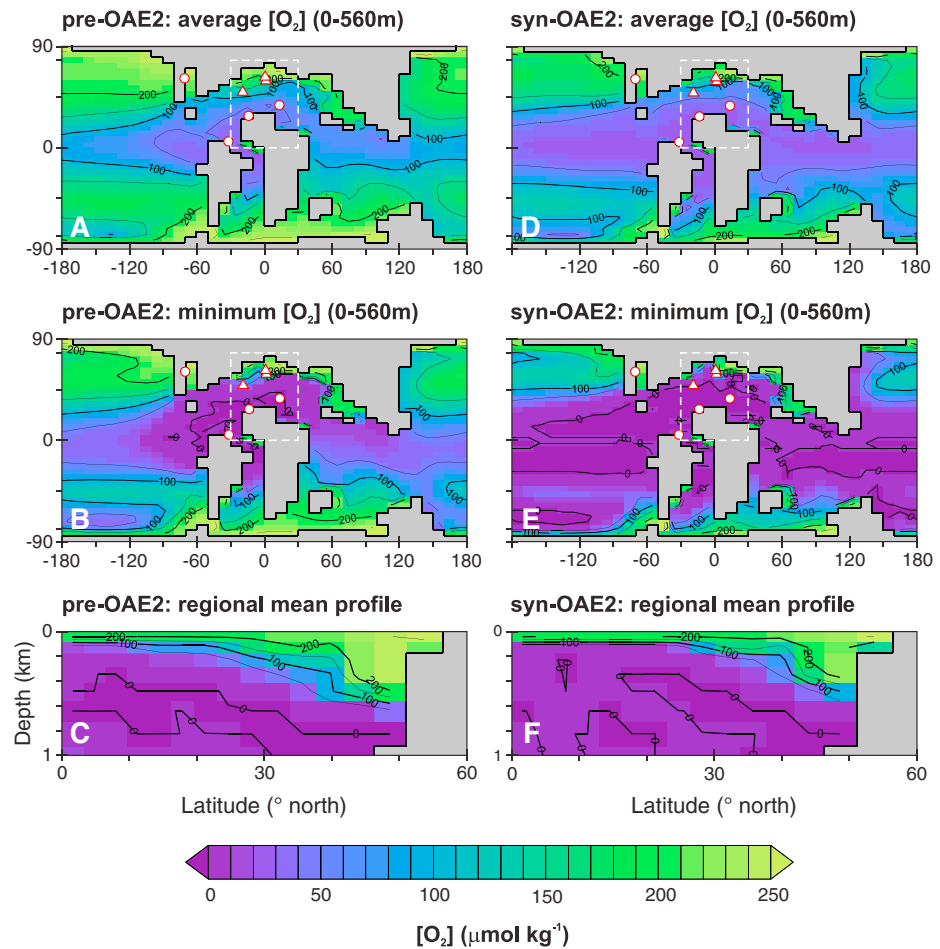


Figure 9. Modeled upper water column-dissolved oxygen distributions (a–c) before and (d–f) during OAE2. Figures 9a and 9d show the averaged O_2 concentration over the upper 560 m of water column; Figures 9b and 9e present the minimum $[O_2]$ in the depth range of 0–560 m. Figures 9c and 9f show the cross-sectional views of dissolved oxygen in the proto-Atlantic upper ocean. The extent of longitudinal averaging in these zonal sections is marked by a white dashed rectangle in the map view panels.

4.5. Data Model Comparisons

To help visualize patterns emerging from the data, the average and ranges of I/Ca are marked on the map in Figure 1 as a simple generalized representation of the overall conditions throughout the studied interval at each site. Three sites located in the NW proto-Atlantic and European shelf sea (Newfoundland Drifts and Eastbourne and South Ferriby, UK) were generally more oxygenated, as indicated by higher background I/Ca , compared with other regions. The area of this oxygen anomaly is small compared to the majority of the proto-Atlantic localities showing lower I/Ca values. By contrast, the paleo-Pacific ocean has been suggested to be largely oxidic [Takashima *et al.*, 2011; Hasegawa *et al.*, 2013], with black shales developed in equatorial regions [Schlanger *et al.*, 1987; Arthur *et al.*, 1988].

Higher dissolved oxygen concentrations are also found in the northeastern proto-Atlantic and European shelf sea prior to OAE 2 in the Earth System Model simulations [Monteiro *et al.*, 2012] (Figure 9). This modeled oxygen maximum occurs only as deep as ~500 m, with the maximum depth decreasing south and away from the continental margin (Figure 9c). Upper water column oxygenation generally decreasing toward the equator is also consistent with proxy data. The I/Ca values are similar for South Ferriby and Eastbourne but are higher than those for Newfoundland Drifts (Figure 1). During OAE 2, oxygen depletion is enhanced throughout much of the ocean in GENIE (Figures 9d–9f). However, the northernmost regions of the proto-Atlantic and European pelagic shelf sea retain an oxygenated water column, consistent with data from South Ferriby showing only very minimal variations in baseline I/Ca during the CIE (Figure 6).

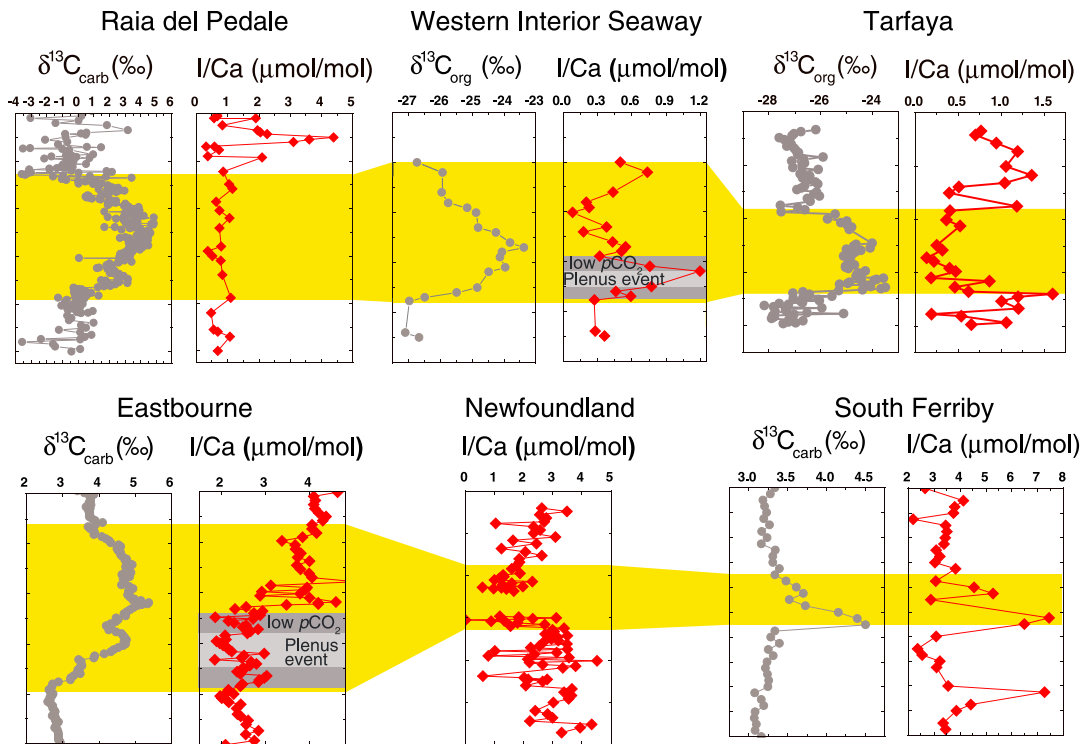


Figure 10. I/Ca records correlated by $\delta^{13}\text{C}$ showing distinct local surface water redox evolution. The yellow boxes mark the CIE, except that the onset is missing at South Ferriby due to a hiatus.

In the model simulations, oxygen depletion is most intense close to the equator on the eastern side of basins due to upwelling fueling biological productivity (and hence subsurface O_2 consumption), as well as subtropical regions of relatively restricted topography. At higher latitudes, the depth of seasonal mixing increases, keeping the upper water column oxygenated. For the northerly proto-Atlantic, mixing results in oxygenation down to depths of up to ~ 500 m, although true intermediate or deep water does not form in the GENIE simulations (unlike the case in the far North Pacific and parts of the Southern Ocean). The proxy interpretation for sites such as South Ferriby is that they were predominantly supplied with oxygenated and iodate-rich surface currents, as opposed to the primarily upwelled and iodate-depleted waters at Demerara Rise and Tarfaya.

The notable exception to the otherwise general model data consistency is the Western Interior Seaway (Figure 1). Epeiric seas are poorly resolved in the cGENIE Earth System Model at this resolution (10° increments in longitude with variations but on average 5° spacing in latitude), posing problems for such a shallow and restricted environment. A second caveat is that, as configured with a highly parameterized 2-D energy-moisture-balance atmospheric component [Edwards and Marsh, 2005], the model fails to exhibit interannual variability and is run to steady state. Hence, the possible role of fluctuating ocean currents and upwelling intensity that might explain redox changes and variability in I/Ca values cannot be addressed. Future modeling work could employ higher-resolution settings and involve simulations of ocean dynamics to improve data-model comparisons for epeiric seas.

4.6. Tempo of Local Redox Changes

During OAEs, the tipping point in the global carbon cycle marked by onset of the CIE does not necessarily coincide precisely with the deposition of black shales [Tsikos et al., 2004] or the buildup of low-oxygen water masses at specific locations. As correlated by the diagnostic carbon-isotope chemostratigraphic profiles, I/Ca records suggest highly dynamic local redox changes that were not always synchronous with the pace of global organic matter burial as recorded by the carbon isotope curves (Figure 10). Demerara Rise is excluded from this compilation due to the presence of diagenetic carbonate in the

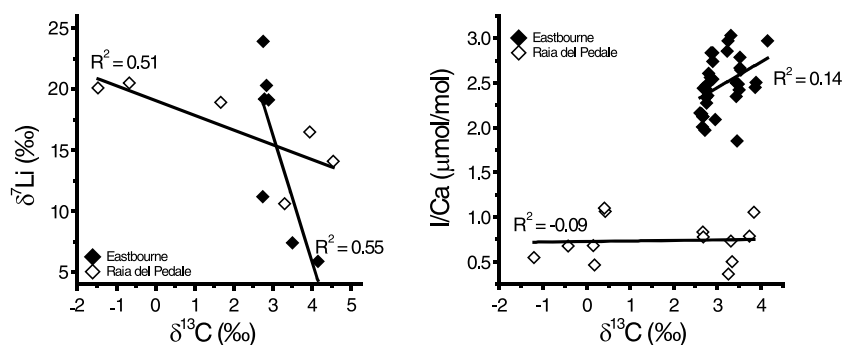


Figure 11. I/Ca and $\delta^7\text{Li}$ plotted against $\delta^{13}\text{C}$ during the rising limb of the CIE for Eastbourne and Raia del Pedale.

core. At two pelagic shelf sea sites (Eastbourne and South Ferriby), I/Ca ratios start to decrease below the onset level of the CIE but not necessarily at the same time, indicating that deoxygenation, likely linked to productivity, began to intensify before the global carbon cycle reached its tipping point and initiated the CIE.

At the Eastbourne site, I/Ca shifted to higher values rapidly and simultaneously with $\delta^{13}\text{C}$ during the middle of the CIE (Figure 10) and remained high during the later stage of the CIE. However, I/Ca reached maximum value at the onset of CIE at the Tarfaya site, decreased rapidly, and shows minimum values stratigraphically just above the depth in core with the highest $\delta^{13}\text{C}$ values. The most reducing local/regional conditions in at least part of the Western Interior Seaway are also registered postdating the peak CIE. I/Ca values during the CIE at Raia del Pedale were consistently low with very little variation and do not vary with changes in carbon isotope values. Similarly, I/Ca continued to rise after the CIE at the Newfoundland Drifts site, Eastbourne, and South Ferriby. This compilation of I/Ca data (Figure 10) indicates that the local oxygen levels in the upper ocean were highly heterogeneous across different sites and different ocean basins during OAE 2, indicating that redox changes did not necessarily follow the global $\delta^{13}\text{C}$ trend.

4.7. Marine Iodine Budget

Globally, the marine iodine budget may be affected by increased continental weathering and organic carbon (iodine) burial. If I/Ca records were predominantly artifacts related only to global changes in total iodine concentration (iodide + iodate), the I/Ca trends at different sites would show considerable similarities, as is observed for the isotopic records of Ca, Li, and C (Figures 2, 6, 8, and 10). Therefore, the differing I/Ca records among the studied sites probably have been affected by local redox conditions more than the potential large-scale changes in total iodine concentration. Sr isotopes ($^{87}\text{Sr}/^{86}\text{Sr}$) increased transiently at the onset of OAE 2, such as at Ocean Drilling Program (ODP) Site 551 in the northeast Atlantic [Bralower *et al.*, 1997] and the southern Apennines of Italy [Frijia and Parente, 2008], suggesting a pulse of enhanced continental weathering, although the overall trend is nonradiogenic. Because Sr isotopes are significantly affected by hydrothermal activity and volcanism, which could overprint the effects of changing weathering rates, this isotopic ratio is not discussed in this study as a weathering proxy [Jenkyns, 2010]. The observed I/Ca trends among sites are very different compared to calcium, lithium, and carbon isotope records. Eastbourne and Raia del Pedale have both Li isotope and I/Ca data covering the rising interval of $\delta^{13}\text{C}$, a period of rapid global organic carbon burial. The two global proxies ($\delta^7\text{Li}$ and $\delta^{13}\text{C}$) show slightly better correlation ($R^2 > 0.50$), compared to I/Ca versus $\delta^{13}\text{C}$ ($R^2 < 0.15$; Figure 11).

Plotting Ca isotopes against $\delta^{13}\text{C}$ and I/Ca separately does not show any major correlation, likely due to the sample resolution and analytical uncertainty in Ca isotope records [Blättler *et al.*, 2011]. Therefore, any global surplus/deficit in the marine iodine budget related to weathering or organic carbon burial was masked by the signal related to dynamic local/regional redox changes in the upper ocean. Although some changes in the iodine budget during global redox shifts and organic carbon burial events are highly likely, this current data set seems to demonstrate that bulk carbonate I/Ca is primarily a proxy reflecting local and proximal upper ocean environmental conditions, at least during OAE 2.

5. Conclusions

Multiple I/Ca records suggest that global upper ocean redox conditions during the Cenomanian-Turonian Oceanic Anoxic Event were remarkably variable. The timing of deoxygenation in the upper ocean was not uniform across ocean basins, and the initial development of poorly oxygenated near-surface waters at some Atlantic locations predated the global carbon isotope excursion. A shallow oxygen oasis in the NW proto-Atlantic and European pelagic shelf sea is supported by both higher I/Ca values in multiple sites and by Earth System Modeling. Comparison between weathering indices and I/Ca records suggests that continental input was not a major driver of changes in the marine iodine budget on the time scale of the OAE. The highly variable stratigraphic I/Ca trends at the investigated sites indicate that I/Ca is primarily a proxy for local redox conditions rather than for global iodine cycling during OAEs. I/Ca and $\delta^{34}\text{S}_{\text{CAS}}$ data both indicate important water mass redox changes, but the latter reflects processes operating on a global scale. This study further demonstrates the potential of I/Ca as a unique paleoceanographic proxy for the identification of relative subtle oxygenation changes in Earth's history [Loope *et al.*, 2013; Hardisty *et al.*, 2014; Zhou *et al.*, 2014].

Acknowledgments

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