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The Cayo del Oso Site (41NU2) Volume II Results of Archeological Monitoring of Spur 3, Corpus Christi, Nueces County, Texas 2000-2007

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The Cayo del Oso Site (41NU2)

Volume II

*Results of Archeological Monitoring of Spur 3,
Corpus Christi, Nueces County, Texas 2000-2007*



by

Barbara M. Meissner, Steve A. Tomka, Jennifer A. Thompson and Raymond P. Mauldin

with contributions by

Vaughn M. Bryant, Linda S. Cummings, J. Philip Dering, Charles D. Frederick, Russell D. Greaves,
Darden Hood, Leonard Kemp, Richard B. Mahoney, Manuel R. Palacios-Fest, Jason B. Perez,
R.A. Varney, Barbara Winsborough and Chad Yost

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Center for Archaeological Research
The University of Texas at San Antonio
Archeological Report, No. 390

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**The Cayo del Oso Site (4INU2) Volume II:
Results of Archeological Monitoring of Spur 3,
Corpus Christi, Nueces County, Texas 2000-2007**

Meissner, Tomka,
Thompson and Mauldin

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Volume II
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County, Texas 2000-2007**

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Abstract:

This report is the second volume in a two volume set (see Jackson et al. 2004) detailing the archeological work conducted at site 41NU2 over several years. This volume focuses on the results of construction monitoring associated with road improvements to Spur 3, in Corpus Christi. The work was conducted under Texas Antiquities Committee Permit Number 2445, originally issued to Dr. Robert Hard, who served as the Principal Investigator during the early phases of the project. In 2001, the permit was transferred to Dr. Raymond P. Mauldin, who served as the Principal Investigator the remainder of the project.

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Chapter 1: Introduction

Barbara M. Meissner and Jennifer A. Thompson

Project Description

In August 2000, archeologists with the Center for Archaeological Research (CAR) of the University of Texas at San Antonio, under contract with the Texas Department of Transportation (TxDOT), began monitoring road, utility improvements, and other construction activities in the right-of-way (ROW) of Spur 3, Ennis Joslin Road, in Corpus Christi, Texas (Figure 1-1). This archeological monitoring took place under Antiquities Permit No. 2445, issued by the Texas Historical Commission (THC), to Dr. Raymond P. Mauldin, Principal Investigator. This is Volume II of a series reporting the results of archeological work at 41NU2 (see Jackson et al. 2004).

The archeological work was necessary to address the requirements of Section 106 of the National Historic Preservation Antiquities Code of 1966, as amended, the

implementing regulations of 36 CFR Part 800, and the Texas Antiquities Code, because the construction impacts land owned by the State of Texas and because the Federal Highways Administration (FHWA) is partially funding the project.

The purpose of the construction was to rebuild and widen Ennis Joslin Road and the eastern-most section of South Alameda Street in order to facilitate transportation from Highway 358 (South Padre Island Drive) to the campus of Texas A&M University—Corpus Christi (TAMU—CC), on Ward Island in the mouth of Oso Bay (Figure 1-2). The road to be rebuilt was known to cross over site 41NU2, the Cayo del Oso site, and the planned road and utility improvements were expected to disturb this prehistoric cemetery and campsite. The site lies on clay dunes along the northwestern shore of Oso Bay and extends from a small creek that empties into the western side of Oso Bay to Corpus Christi Bay, approximately 960 m to the north (Figure 1-2). The width of the site is known only

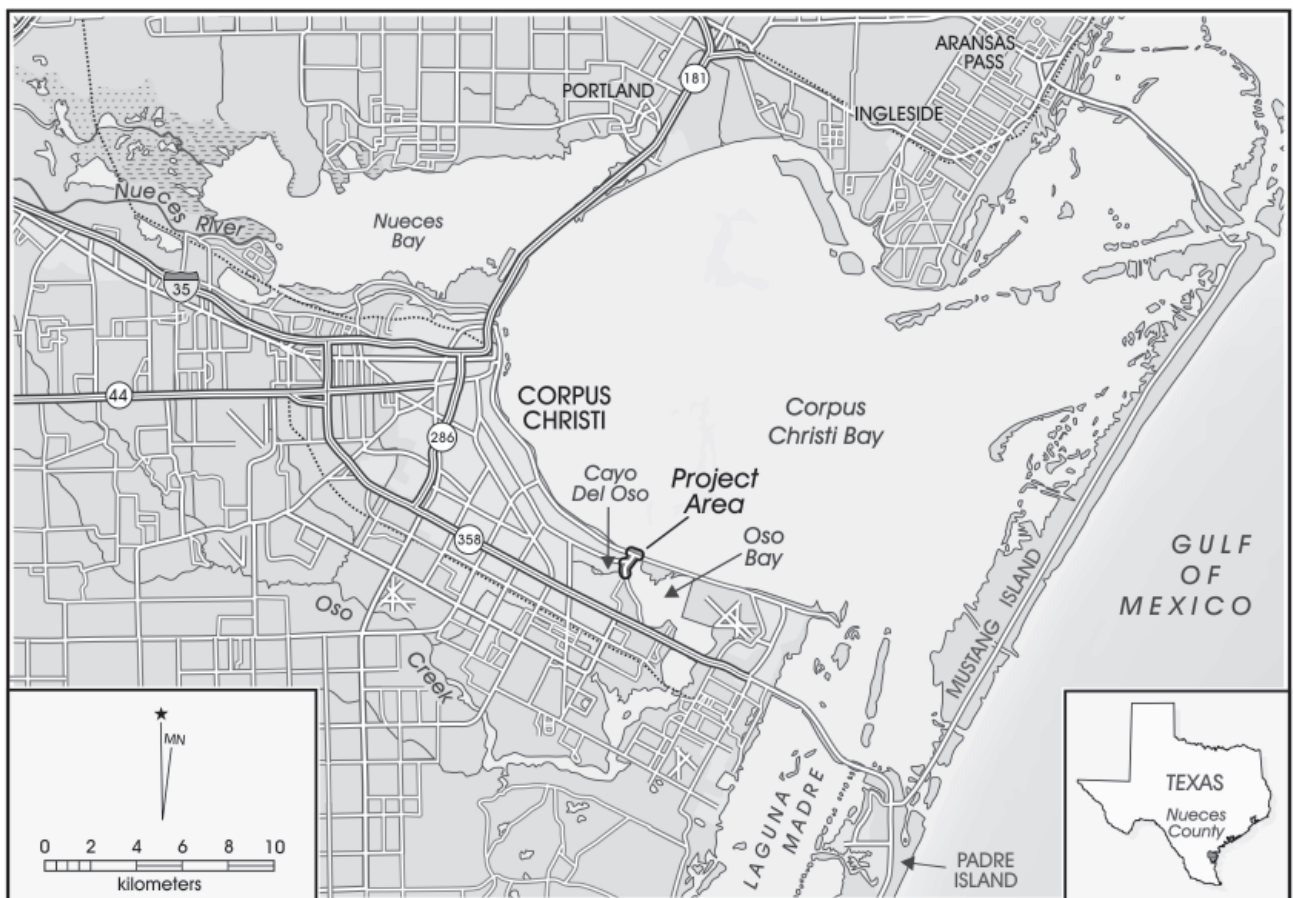


Figure 1-1. Map of the Corpus Christi Bay system, including Nueces Bay, Oso Bay and the northern Laguna Madre.



Figure 1-2. A map of the Cayo del Oso site superimposed on an aerial photograph taken in 2004.

on its northern border, on a bluff overlooking Corpus Christi Bay. At that point, the site is approximately 200 m wide. The area is a busy urban residential neighborhood, with a small commercial area along Ocean Drive.

The general locale has been known by a number of different names, including Callo del Oso, Calle del Oso, Cale del Oso and Cayo del Oso. *Cayo* means key (i.e., low, coastal island) in Spanish and is probably the correct name, as there are many small, ephemeral, and constantly changing islets all over the mudflats, and Ward Island (see Figure 1-2) could definitely be classified as a “key.” Even though most of the previous literature about the site refers to it as Callo del Oso (Martin 1930; Meissner 2000, 2003; Ricklis 1997), probably following Martin (1930), the area, also known as the False or Blind Oso (Figure 1-2), is referred to as Cayo del Oso in government documents (see General Land Office 2008), including the TxDOT construction maps for this project. Beginning with Volume I of this series, therefore, we have chosen to refer to the site as Cayo del Oso.

The site has been known as a prehistoric cemetery and campground since at least the last quarter of the nineteenth century. For instance, in a footnote to a paper published in 1898, there is a reference to boyhood memories of the area by an acquaintance of the author, who reported “[human] bones on the Cala del Oso have been found in quantities” (Coopwood 1898:117).

Volume I of this series (Jackson et al. 2004) provides a detailed history of what is known about excavations by various amateur and professional archeologists in the 1920s to the 1960s. These investigations are briefly summarized in the last section of this chapter. It is known that many of the burials excavated from the site were scattered to various locations and their whereabouts are no longer known (Jackson et al. 2004:17). Remains of at least 110 individuals that were excavated before 1970 are curated at the Texas Archaeological Research Laboratory (TARL), and it is known that there were many more excavated at the site. Of the 110 burials curated at TARL, 105 came from the excavations led by A. T. Jackson

of the University of Texas at Austin, under the auspices of the Works Progress Administration (WPA) during the summer of 1933. Volume I of this series published, for the first time, Jackson's draft report of his excavations, on file for more than 70 years at TARL (Jackson et al. 2004).

As mentioned above, the location of the site has been known for a century or more. The site has been actively excavated since at least the late 1920s (Baird 2006; Martin 1930) and possibly longer. Examination of field notes and photographs from excavations at the site in the 1930s indicated that there had been extensive erosion of the clay dunes, and it was known that the site had been subjected to many years of excavation by both amateur and professional archeologists, and many casual diggers. Because of this, it was considered unlikely that the site still contained undisturbed sediments with substantial research potential (e. g. Hall 1995a:49). However, the number of burials recovered at the site in the past made it important to test the project ROW within the site boundaries to determine how likely it was that the project would disturb human burials or important cultural materials. Therefore, in 1996 TxDOT contracted with Coastal Archaeological Research, Inc. to test selected areas of the ROW (Ricklis 1997). The testing encountered one human burial in the southern part of the site, as well as many areas where undisturbed sediments contained cultural materials in a series of sealed strata (Ricklis 1997:30, 33, 72-73). Ricklis' (1997) results will be discussed in more detail in the last section of this chapter. As a result of the testing, Ricklis (1997:72-73) recommended that all construction within the site boundaries be monitored by professional archeologists. TxDOT then contracted with CAR to provide this service.

The purpose of the monitoring was to identify any cultural remains that may be within areas of impact and to "... complete the documentation of features and the exhumation of any human burials as quickly as possible to allow the construction contractor to regain access to the area" (Exhibit 1, WA #573 11SA002). With intermittent interruptions, construction monitoring continued from August 2000 until February 2006.

This volume reports the results of the construction monitoring, as well as a pilot study in Human Behavioral Ecology (HBE), which focuses on the material recovered from 41NU2. The document is organized into fourteen chapters and seven appendices. The remainder of this chapter includes a very brief presentation of the nature of the research design developed for this report, and a summary of archeological work that has occurred at 41NU2.

Following this introductory chapter, Chapters 2 and 3 provide information on the geology, environmental

structure, and cultural history of the Corpus Christi area. Chapter 4 contains a detailed presentation of the field and laboratory methods employed. Chapter 5 summarizes the work accomplished, the human remains identified, and the artifacts recovered during monitoring. Chapter 6 discusses the correspondence between radiocarbon dates obtained on charcoal and shell from the site, the ages of the archeological components, and the features and occupation zones defined. Chapter 7 lays out the theoretical framework and Chapter 8 describes the regional resource structure. Chapter 9 presents a model of hunter-gatherer adaptation and regional and site-specific expectations. Chapters 10-13 contain discussions of paleoenvironmental reconstructions (Chapter 10), subsistence (Chapter 11), land use (Chapter 12), and mortuary practices (Chapter 13). Chapter 14 presents the conclusions, bringing the information in the previous chapters to bear on the HBE model to determine its usefulness in future research projects on the Texas Coast. Appendices A through G are the reports provided by experts in the fields of pollen and phytoliths (Appendix A and B); macrobotanical analysis (Appendix C); diatom analysis (Appendix D); ostracod, foraminifer, and micro molluscs analyses (Appendix E); several geological analyses (Appendix F); and radiocarbon sample data forms (G).

Research Design

The research design, developed as a collaborative effort between CAR and the TxDOT Environmental Affairs Division, focuses on the interpretation of data made available by the monitoring project at Spur 3 (Tomka et al. 2006). Using principles derived from the habitat choice model (Sutherland 1996), we developed a general model of coastal hunter-gatherer settlement pattern and subsistence practices. Next, we defined a series of site-specific expectations related to the function and use of 41NU2 and how the nature of site use may have changed over time. As a test of these expectations, we proposed to study several data types recovered from 41NU2 and compare trends within them to data types from a small number of previously excavated sites reported in the regional literature. The pilot study served as a test case of the usefulness HBE approaches as a research tool to model hunter-gatherer behavior.

The research goals described in the research design (Tomka et al. 2006) rely on the assumption that the materials observed and collected by CAR from 41NU2 cannot be interpreted if viewed in isolation. This perspective is especially germane to sites located along the Gulf Coastal strip since ethnohistoric and Late Prehistoric settlement pattern reconstructions (see Ricklis 1997, 2004, Ricklis and Weinstein 2005) indicate a

strong reliance by hunter-gatherer groups on both coastal and inland zones.

To provide the needed perspective, detailed data was collected from site reports on 18 sites near 41NU2 for use in the comparative studies described in this volume. In addition, less inclusive data on all prehistoric sites in Nueces County and three adjacent counties, Aransas, Refugio, and San Patricio, was collected from THC's online Texas Archeological Site Atlas (THC 2008). This data is limited by the variability of information included on Site Report forms in more than forty years of site data recording stored at TARL. To the extent possible, given these limitations, this data was used in GIS-based studies of prehistoric settlement patterns across the four counties. A more detailed GIS-based study of prehistoric cemeteries in South Texas has also been completed and the results are presented in Chapter 13.

Previous Excavations at 41NU2

Volume I of this series (Jackson et al. 2004) describes early excavations at 41NU2 between 1900 and 1970 and includes a manuscript of A. T. Jackson's report of his work at the site. This section of Chapter 1 provides a very brief review of the work conducted at the site prior to 1970 (Jackson et al. 2004), followed by brief summaries of the archeological investigations done by Coastal Archaeological Research, CAR, and SWCA Environmental Consultants associated with the Spur 3 construction project.

Before 1970

As mentioned above, a detailed discussion of what is known about excavations at 41NU2 between about 1928 and 1970, including the first publication of Jackson's draft report on his 1933 excavations, is available in Volume I of this series (Jackson et al. 2004).

While the site was known at least since the late 1890s (Coopwood 1898), by the late 1920s rumors of the site had reached the ears of George C. Martin, an avid amateur archeologist who had, with his friend Wendell H. Potter, just completed a massive survey of the central upper southern Texas coast (Martin and Potter n.d.). Though a few copies of their privately printed report still exist, no date of publication is available.

During exploration of the site in 1929, Martin (1930) was interviewed John B. Dunn, who told him he had worked on the property during the Civil War and that the land was then heavily vegetated in brush. Mr. Dunn claimed there was no

sign of a prehistoric site there at the time. Martin (1930) while excavating at the site, was told by the land owner that after the great hurricane of 1900, he returned to his property and saw human bones eroding out of the clay dunes in numbers he described as "over 5000" individuals (Martin 1930:9). While it is easy to suppose that this number is a gross exaggeration, it is clear that the man had seen a large number of bones he believed to be human. Mr. Dunn also indicated that the hurricane had heavily eroded the site, exposing burials and other cultural materials (Martin 1930:9). Martin recorded finding a total of 53 skulls during several trips to the site (Jackson et al. 2004).

At some time after Martin's excavation (probably in late 1931), E.B. Sayles excavated a single test pit and recovered 9 more burials (Sayles n.d.). His results and Martin's (1930) report encouraged A. T. Jackson, of the University of Texas, to put together a field season at the site, under the auspices of the WPA, during the summer of 1933. He excavated an area known as the "bone pit", recovering what he believed at the time to be 93 individuals. Later investigations at the site consisted of small-scale excavations by such professional archeologists and anthropologists as Alex Kreiger and T. N. Campbell, as well as such scientists as W. A. Price, a geologist who published descriptions of how clay dunes accumulate (Huffman and Price 1949; Price and Kornicker 1961) and others who returned their findings to TARL. In addition, a large number of individuals, amateur archeologists and collectors, undoubtedly also dug at the site (Jackson et al. 2004).

Archeological Projects Associated with the Spur 3 TxDOT Project

When Corpus Christi State University, located on Ward Island, was transferred to the Texas A & M University System in 1989, and renamed in 1993 as Texas A & M University—Corpus Christi, a plan was developed to enlarge the campus (TAMU—CC 1989). As part of this plan, the City of Corpus Christi and TxDOT noted that Ennis Joslin Road, the street most students and faculty used to get to campus, was inadequate in width for the increased volume of traffic to TAMU—CC. The road had been originally planned as a small residential street and by the mid-1990s was deteriorating rapidly under the pressure of the increasing volume of traffic. Plans were made to widen Ennis Joslin Road and improve its intersections with various side roads and Ocean Drive (see Figure 1-2). As part of these road improvements, the City of Corpus Christi would take the opportunity to reorganize and upgrade the city underground utilities along the road (R. Moff, personal communication 2000). Recognizing that this might adversely impact 41NU2, in 1996 TxDOT began to

explore the condition of the site within the proposed project area, the Area of Potential Effect (APE).

Coastal Archaeological Research, Inc.

Coastal Archaeological Research, Inc. was contracted to conduct testing within the archeological APE. In 1996, Ricklis (1997:16) excavated a series of thirty 2-inch cores and twenty test units. The test units were dug with a Gradall excavator and were roughly 1.5 by 2.2 m (see Ricklis 1997:25 for locations of these units). The units exposed three major soil profiles, all underlain by the culturally sterile Pleistocene Beaumont Clay (see Chapter 2). These were: 1) disturbed clay loam over a gray brown sandy clay characterized as a “upland prairie soil profile” (Ricklis 1997:29); 2) a series of lenses of light and dark sandy clays consistent with clay dune formation; and 3) a black homogenous highly organic clay described as pond sediment, representing a Late Holocene ephemeral pond that developed behind a clay dune (Ricklis 1997).

Ricklis (1997) reported cultural material from previously undisturbed strata in fourteen of the units. Cultural deposits included scatters and concentrations of estuarine shells, charcoal, burned clay nodules, a few lithics, shell and bone tools, and faunal bone within the darker bands of the clay dune and directly above the Beaumont Clay in the non-dune center of the site. In Unit 5, located on the east side of Ennis Joslin Road in the southern part of the project area, a single relatively intact human burial was encountered (Ricklis 1997:23, 29).

Though he noted the thinness of the cultural deposits he encountered and did not believe that the camp debris would provide important archeological information if excavated, Ricklis’ recorded six features in addition to the burial (Ricklis 1997:35-41) His recommendations (Ricklis 1997: 73-73) included the need for archeological monitoring of construction at the site within the TxDOT ROW. He also noted that Right-of-Access issues had limited the areas he had tested, leaving unknown the likelihood of undisturbed cultural deposits and/or human burials in those areas (Ricklis 1997:3, 73).

Center for Archaeological Research

In 2000, under contract with TxDOT, CAR began monitoring the construction excavations for exposure of important archeological components and human interments and reported findings in a series of interim reports submitted to TxDOT (Meissner 2000, 2003, 2004, 2005). Between August 15 and October 3, 2000 a total of 16 features, consisting of concentrations of shell, charcoal and vertebrate faunal remains were excavated, screened, and all recovered cultural

materials collected, as required by TxDOT (see Chapters 4 and 5). On October 3, 2000 TxDOT decided to stop the project temporarily while considering the unexpected number of utilities that were not on any city utility maps, as well as the possibility that construction plans would be changed.

The original plans for details of construction, including utility placement, and scheduling of each component of the construction were changed. While this was being done, construction activities at 41NU2 were sporadic, consisting of digging to locate and/or move underground utilities and grading of an area that would be temporarily paved as a bypass during later construction (Meissner 2003). Whenever construction took place on the site, TxDOT personnel informed CAR and archeologists were present to monitor activities.

Beginning in April 2003 and continuing to November 2005, with occasional brief interruptions, CAR monitored the remainder of the project. A total of 20 major utility trenches, as well as 14 lateral trenches that crossed the road and numerous shallow trenches to attach homes and business to the new system, were completed. In addition, several large pits were dug to place manhole boxes. Ten large areas where sheet midden deposits were seen in profile were identified as Occupation Zones (Meissner 2004, 2005). A total of 170 features, including the 16 found in 2000, were identified in the field. However, beginning in April 2003, TxDOT required that features be minimally recorded (see Chapter 4) and stated that locating human remains before they were disturbed by construction activities, should be the major focus of CAR’s monitoring.

Six human burials were encountered during the monitoring. Only Burial 2, the first (as the burial located by Ricklis (1997:42) was considered Burial 1) located, was completely intact. Burials 3 to 5 were incomplete, being disarticulate bones found in the trench fill of old utility trenches and presumably disturbed and reburied during the earlier utility construction. Burials 6 and 7 were the partially disturbed remains of a child and an adult. Though the condition of the bones and the degree of disturbance made it difficult to be sure, it appeared that the adult burial had been disturbed by the burial of the child. As one of the deliverables under this project, in addition to a standard osteometric analysis, the CAR completed the NAGPRA human remains inventory. This inventory and the results of the analysis were submitted to TxDOT-ENV in 2007 (Mahoney and Shafer 2003).

In February 2006, CAR monitors returned to observe the planting of palm trees in previously undisturbed areas. No new features were identified. In August 2007, CAR archeologists returned with the geomorphologist Charles

Frederick to bore a total of six cores in order to obtain soil samples for various specialized analyses, including lithology, pollen, phytolith, ostracod and diatom studies. Results of these studies are incorporated into this report and can be found in the Appendices.

SWCA Environmental Consultants

In March and April 2006, SWCA Environmental Consultants, under SWCA's blanket Texas Antiquities permit #3740, and with Kevin A. Miller acting as Principal Investigator,

monitored the excavation of three electrical conduit poles and 48 light pole footers (Miller et al. 2006:3). The majority of excavations monitored by SWCA were in areas that had previously been disturbed by the placement of utilities during the period of the project monitored by CAR. Three features (all defined as "non-significant" under the terms of the instructions given by TxDOT, see Chapter 4) were identified, each consisting of a concentration of shells, animal bone and charcoal. As per the TxDOT instructions, no samples of feature fill were taken (Miller et al. 2006:14-15). The location of the excavations and the three features were mapped a (see Chapter 6).

Chapter 2: Background, Geology, Climate and Environmental Structure

Steve A. Tomka and Raymond P. Mauldin

This chapter briefly summarizes the geology of the Corpus Christi Bay system. The summary is followed by a discussion of the modern climate of the area. The environmental structure that results from the geology and climate of the area is then discussed.

Development of the Corpus Christi Bay System

Geological History

During the last glacial maximum ca. 20,000 B.P., the area that is now Corpus Christi was located on uplands near the deeply incised proto-Nueces River valley. The coast was approximately 200 kilometers east-southeast of the current coastline, as sea levels were about 100 m lower than modern levels (Brown et al. 1976; Ricklis 1995a:18). During an earlier interglacial, a thick layer of clays and sandy clays, known as the Beaumont Clay Formation, was deposited as a result of fluvial and colluvial processes that created a deltaic plain across most of what is now the Gulf Coast of Texas (Brown et al. 1976:20). The Pleistocene Nueces River deeply incised into this formation as the sea level dropped during the most recent glacial maximum.

When the most recent interglacial began ca. 18,000 B.P., ice melted and the sea level began to rise. The details of the rate at which this occurred are still being researched (Ricklis 1997:6), however, there is a consensus that sea level rise was rapid beginning at 18,000 B.P. until about 9000 B.P., after which the rise became more intermittent. Each time there was a significant rise in sea level, estuary systems formed since the previous sea level rise were inundated, decreasing their potential for production, until a new estuarine environment could develop at the higher level (Ricklis 1997:6).

The inundation of the Pleistocene Nueces River Valley near Corpus Christi started about 9000 B.P. with the incursion of marine water up the channel of the river in its deep valley as far as 13 km north of the current mouth. The Nueces River began a headward delta formation that was later inundated as the sea level rose again. At this time, about 7500-7000 B.P., the entire valley was inundated, creating a narrow, deep bay (Brown et al. 1976:20-21). During the next few thousand years wave action and storms eroded the edges of the bay, widening it except at the point where today Nueces Bay feeds into Corpus Christi Bay through a fairly narrow channel (see Figure 1-1). New erosion continued filling in the deep Pleistocene valley, as the Nueces River and Oso Creek

gradually prograded headward, creating the deltas through which they now enter the bays.

Some postulate a highstand as much as one meter above the current sea level from about 5300 to 2800 B.P. (Paine 1991; Ricklis 1997:7). If this highstand did, indeed, occur, it could explain the formation of the False Oso mudflats, as the northwestern side of this larger and wider Oso Bay would have been more susceptible to wave action, as it was downwind of the prevailing winds (c.f. Brown et al. 1976:22). When the waters receded to the current level, the western side of Ward Island had accumulated more sediment than the eastern and was no longer an outlet into Corpus Christi Bay except during storm events.

The development of the modern Corpus Christi Bay environment began when roughly modern sea level was established allowing the development of new shallow estuarine environment. At the same time, sands and clays off the coast began to form small shoals about 5,000 years B.P., developing into a series of small, narrow islands (Fisk 1959). There is some difference in the estimates of when these islands coalesced into roughly the shape and extent of the modern barrier islands of the coast of Texas. Estimates range from roughly 4000 B.P. (Fisk 1959:123) to 2800 B.P. (Brown et al. 1976:21), to between 2,500 and 2,000 years ago (Morton and Paine 1984:12). Similarly, Ricklis (1995a), based on his research along the Nueces Bay shoreline, concludes that the isolated barrier islands coalesced sometime after 3000 B.P. giving rise to the estuarine conditions characteristic of the Nueces and Corpus Christi Bay region today (Ricklis 1995a:142). In any case, open marine conditions in Corpus Christi Oso Bay ended as tidal inlets closed. As the number and depth of passes between the barrier islands decreased, the freshwater entering Corpus Christi Bay had only two or three narrow passages out to the sea. In general, the salinity of the bays decreased, although in periods of extreme drought the salinity in the bay can be as much as 45ppm (Brown et al. 1976). At the same time, the erosion of the bay edges made the bay more shallow, and large areas of estuarine marshes and sea grass flats (see below) developed, creating an ideal environment for many plant and animal communities, as well as nursery grounds for many marine fishes (see Chapter 8).

The Modern Corpus Christi Bay Estuary

Today Corpus Christi Bay is a large, shallow bay that covers a little more than 39,000 hectares (390 km²) in surface area, not including Nueces Bay and Oso Bay (Anderson 1960),

and has a maximum depth of 4 m (Harwood 1980). Corpus Christi Bay receives freshwater inflows from the Nueces River and Oso Creek. The bay is connected to the Gulf of Mexico through two channels: the largest of these is Aransas Pass, a ship channel; the second is Fish Pass dredged near the southern end of Mustang Island (Morton and Pieper 1977). Diurnal tidal fluctuations within the bay have mean amplitudes of 10 cm while astronomical tidal fluctuations range from 0-60 cm. Wind direction and velocity influence the duration and amplitude of inundation of exposed tidal flats.

Oso Bay is an enclosed, soft-sediment, estuarine secondary bay located on the southwestern shore of Corpus Christi Bay. The only major freshwater inflow it receives is from Oso Creek. The average depth of Oso Bay is less than 1 m. Tidal exchange with Corpus Christi Bay occurs through a small pass located east of Ward Island. The wind-tide flat west of Ward Island is approximately 28 hectares (Figure 1-2). This flat is known as the False Oso or Blind Oso, and it is on the western shore of the False Oso that site 41NU2 is located (Figure 1-2).

The mudflats of the False Oso slope gradually under water to the south, where the bay remains no more than a few centimeters deep for several hundred meters. The False Oso mudflats are regularly submerged for a few hours during unusually high tide events, particularly when these high tides

are combined with strong southeasterly winds or tropical storms, or when there is a large influx of fresh water from Oso Creek during major inland rain events (Figure 2-1). The mud flats of the False Oso are the source of the clay dunes that enclose the Cayo del Oso site (Price and Kornicker 1961).

Clay Dune Formation

As mentioned above, the Cayo del Oso site is located on the northwestern shore of Oso Bay, adjacent to a wind-tide flat, on two clay dunes resting on top of the Pleistocene Beaumont Clay Formation. Similar clay dunes are a feature of the Texas coast from Copano Bay to the Rio Grande delta (Huffman and Price 1949:118). They are always found adjacent to the lee of wind-tide flats that are only occasionally inundated. As these flats dry after a rain storm or an inundation, they form a thin crust, consisting of particles of clay and/or silt and salt crystals bound together by algae. As this crust dries, it tends to crack, exposing a more friable layer. Under the pressure of almost constant prevailing winds sand-sized particles of clay, salt crystals and algae form, which are susceptible to aeolian uptake and redeposit onto the shore, over time developing into a clay dune (Huffman and Price 1949:120). The amount of clay deposited varies a great deal, depending on season, weather, and climatic conditions. Small amounts of clay may accumulate during the dry months of summer. During droughts, when vegetation on the flats dies, large amounts of



Figure 2-1. *The False Oso mudflats as they usually appear, taken from the northern end of the Cayo del Oso site, facing east. Ward Island and TAMU-CC are in the background.*

clay particles are blown by prevailing winds onto the surface of the dune (Huffman and Price 1949:120).

Unlike sand dunes, the clay particles do not continue to be susceptible to movement by subsequent winds, as the clay particles “melt” into the compacted clay of the surface of the dune when any moisture is present. Even the dew of the morning is enough to complete this process (Huffman and Price 1949:124). Clay dunes are susceptible to erosion through runoff, especially after a drought when vegetation on the surface of the dune is sparse. The rains that end the drought often cut deeply into the surface of the dune. These erosional features are then gradually filled in. Hurricanes and other big storms can cause a great deal of erosion to clay dunes while at the same time depositing coarse-grained materials.

During periods when there is sufficient rain, the dunes stabilize and vegetation grows thickly on the surface leading to reduced sheet erosion. Over time a much darker, more organic, soil is formed, until it is buried by rapid accumulation of clay during long droughts.

Climate

This section provides a summary of the present climate of the Texas coastal region. Here we define the coastal and near-coastal setting as a roughly 125-km-wide strip of land following the current gulf coastline, roughly including both the Coastal Prairies and Interior Coastal Plains sub-regions of the Gulf Coastal Plains.

Figure 2-2 indicates the area we have defined as the coastal and near-coastal region, an area we will refer to simply as coastal. Also identified in the figure are the locations of 53 weather stations that provide data for the discussion of the modern climate of the area. These data are based on 30 year normal precipitation and temperature records from 1971 through 2000, and come from a data base with nearly 600 stations from across Texas (NCDC 2004).

Precipitation

Figure 2-3 presents the average monthly rainfall for the 53 coastal locations shown in Figure 2-2, along with rainfall patterns from 53 randomly selected sites from the remaining 544 weather stations from across the state. Comparisons of the

two curves clearly suggest that while both data sets reflect a bimodal pattern, coastal stations have different peaks, with most rainfall occurring in September, and a secondary peak occurring in June. In contrast, the 53 stations from the remainder of the state have highest rainfall totals in May, with a lower, secondary peak in October. In addition, in this non-coastal data set the lowest rainfall totals occur, on average, in January, while February and March are the lowest totals for the coastal locations. Finally, note that in every month of the year, coastal sites exceed the average rainfall of more inland settings. At a yearly time scale, the average rainfall for the coastal sites (40.56 in.) exceeds the non-coastal sample average (32.57 in.) by about 8 in., with most of that difference occurring over the mid-summer and early fall months of July through October.

An examination of precipitation data from the 53 coastal stations suggests that there is considerable variability in the annual rainfall totals, from a high of 62.75 in. near the Louisiana border, to a low of 21.69 in. near the southern coastal border with Mexico. In order to explore that variability, we partitioned the data set into three groups by location, with the

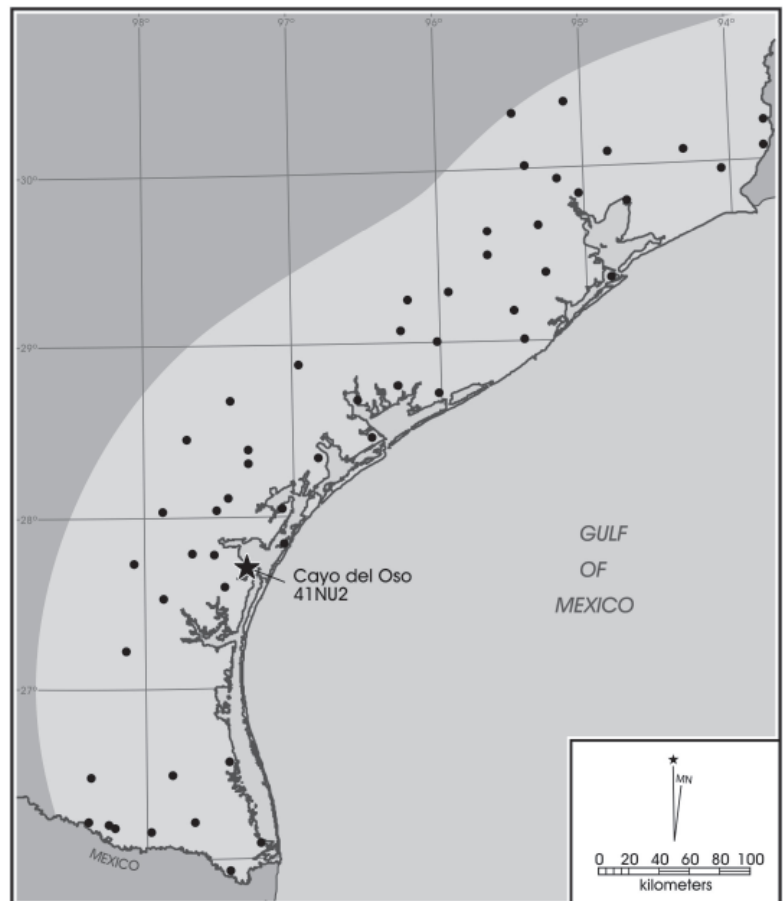


Figure 2-2 Weather station locations within the coastal and near coastal region along the Texas Gulf Coast (NCDC 2004).

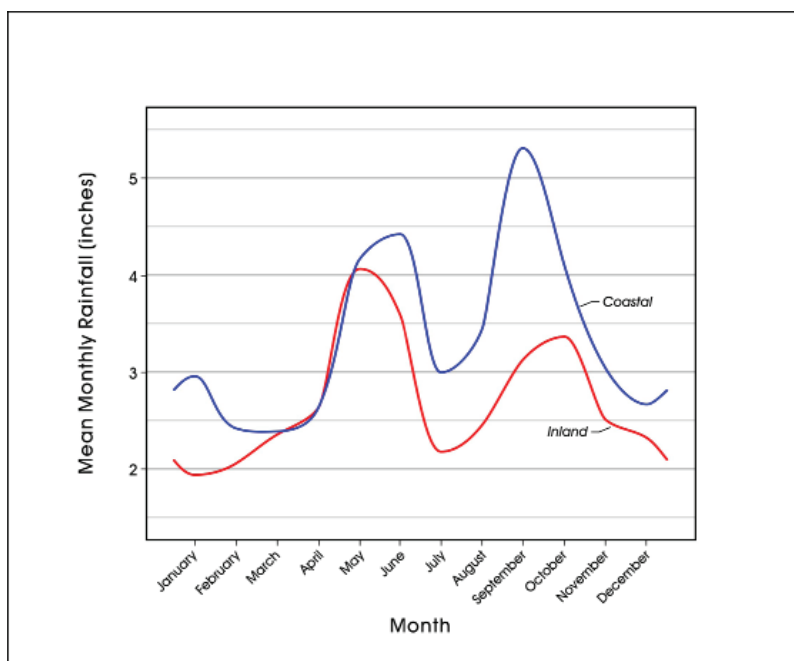


Figure 2-3. The mean monthly rainfall for the 53 coastal locations compared to the average monthly rainfall from the remaining 544 weather stations in Texas. Based on data from NCDC (2004).

first group consisting of 13 stations located in the Southern or Lower section of the coast, a second group of 22 stations in the Middle Coastal region, and a group of 18 stations located in the Upper Coastal area. Figure 2-4 presents these locations (insert), along with their respective monthly average totals between 1971 and 2000. Examination of the figure will show that there is a south to north trend in the rainfall data, with the Upper Coast having significantly higher rainfall and the Lower Coast having significantly lower rainfall. At a yearly scale, the Lower Coastal stations average 25.78 in., with the Middle Coast averaging 39.01 in., and the Upper Coast averaging 53.12 in. The Figure 2-4 line graph patterns suggest that while September peaks are roughly comparable in all three coastal sections, there is significantly more precipitation in the remaining months of the year as one proceeds up the Texas coastline. In fact, the Lower Coast is, on average, drier than the 53 randomly selected stations from the remaining portion of the state.

Figure 2-5 presents mean monthly rainfall totals from 1971 through 2000 for Corpus Christi, located at the southern end of our Middle Coastal region. Overlain on that plot is the monthly averages for the Middle Coast (see Figure 2-4). While following the general shape of the rest of

the Middle Coast, monthly rainfall is lower at Corpus Christi than the average for the Middle Coast region, with much of this deficit coming in the late fall, winter, and early spring months, between November and March.

Kier and White (1978) present 42 years of monthly data for Corpus Christi between 1932 and 1973 which allows an assessment of year to year variability (Figure 2-6). In these data, annual rainfall totals varied from a low of 14.66 in. (1963) to a high of 44.25 in. (1960), with an average yearly amount of 28.71 in. over these 42 years. The year to year variability is substantial, a factor also shown by monthly differences. The highest monthly total during these 42 years was 20.33 in. of precipitation, recorded in September of 1967. While high totals are frequently associated with the summer months of August, September, and October, probably as a result on late summer gulf storms, high totals occur in several other months, including 10.78 in. of rainfall during January 1958 (Kier and White 1978).

Temperatures

Figure 2-7 presents box plots of average annual temperatures (°F) for the 53 coastal stations as well as 53 non-coastal locations selected at random from the remaining portions of

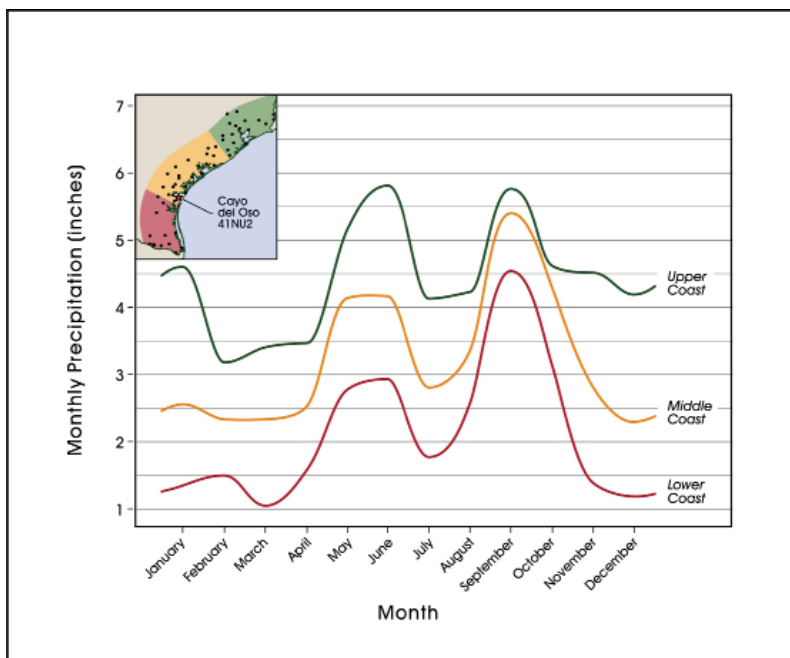


Figure 2-4. The mean monthly rainfall for the Upper, Middle, and Lower coastal regions. The inset defines these regions. Based on data from NCDC (2004).

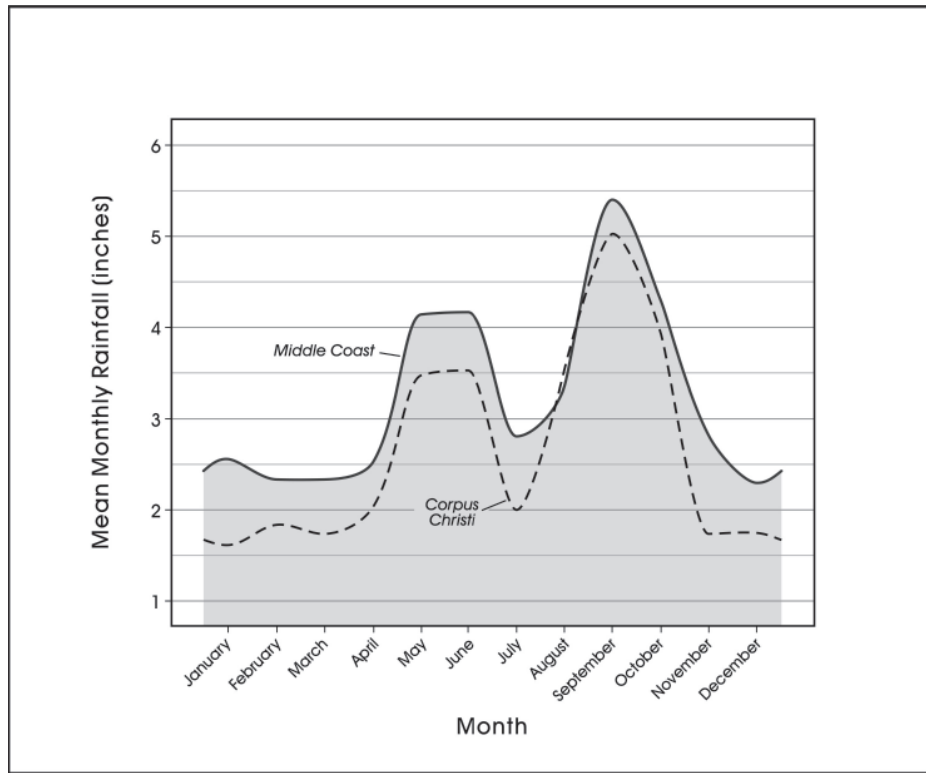


Figure 2-5. Mean monthly rainfall totals for the Middle Texas Coast and Corpus Christi (1971-2000). Based on data from NCDC (2004).

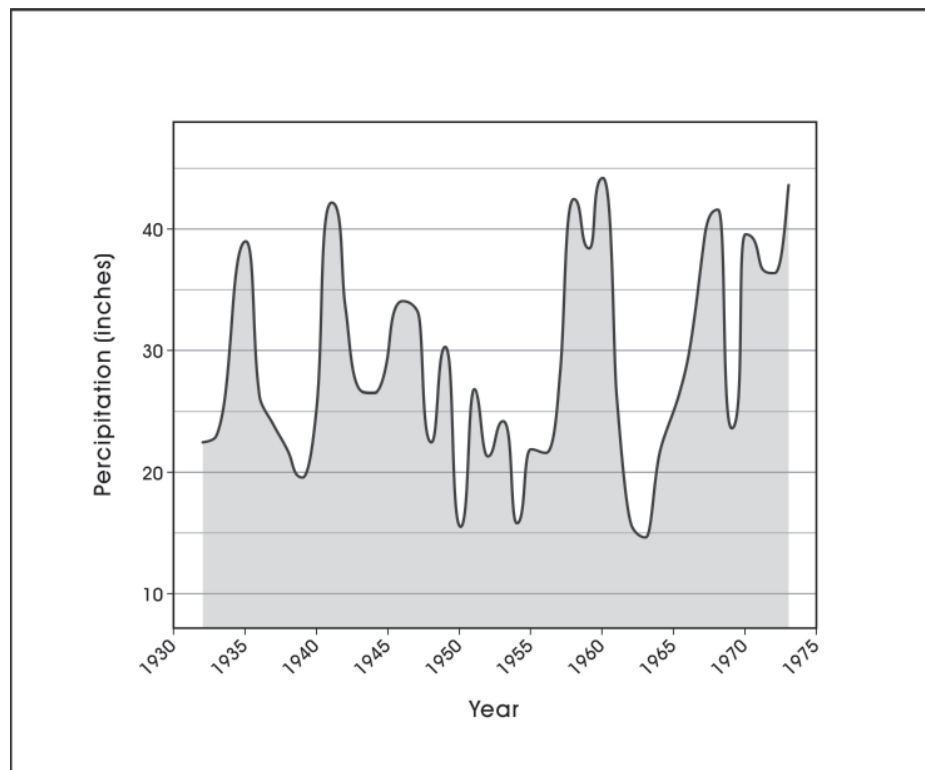


Figure 2-6. Annual rainfall at Corpus Christi between 1932-1973. Based on data from Kier and White (1978).

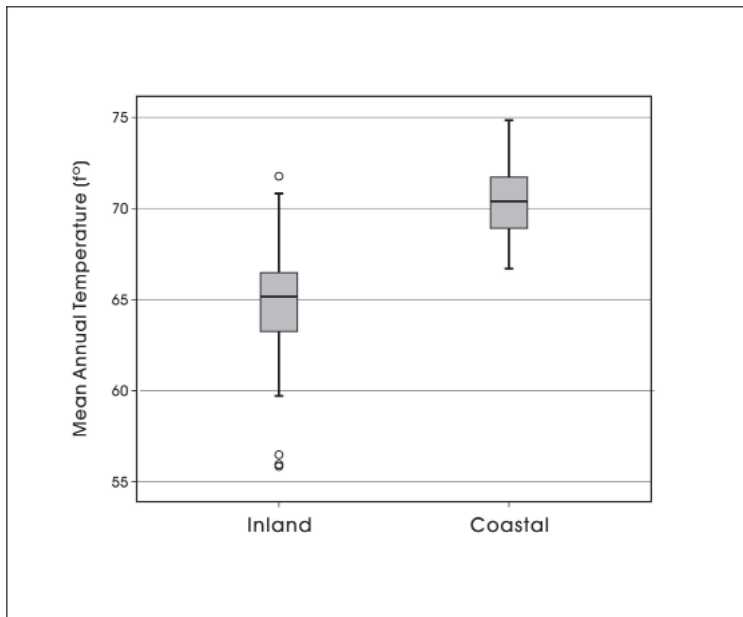


Figure 2-7. Mean annual temperatures of coastal weather stations compared to inland weather stations. Based on data from NCDC (2004).

Corpus Christi Bay water temperatures vary from a low of 7°C in the winter to 29°C in the summer (Harwood 1980).

Salinity

Average annual salinity for the bay is around 30 parts per thousand (ppt). In periods of high rainfall and runoff, salinity can drop to as low as 11 ppt, while periods of low rainfall and high temperatures may increase salinity to 45 ppt (Brown et al. 1976). Figure 2-10 shows the clear relationship between rainfall and average salinity in Oso Bay (Brown et al. 1976). The same pattern also should replicate on a seasonal basis in Corpus Christi Bay proper, although the amplitude of the effects should be less due to the size and depth of the larger bay.

Southeasterly winds are typical of the area throughout much of the year. However, strong northerly fronts are common in the winter and can result in below-freezing

the state. Clearly, coastal settings are significantly warmer during the 1971 to 2000 period, with mean yearly temperatures for the coastal group exceeding the random samples average by 5.65°F. As with rainfall, there are also significant temperature differences between our three geographical distinctions as can be seen in the box plots in Figure 2-8. Average January (Top) and July (Bottom) temperatures are significantly higher in the lower coastal area, while the upper coast has the lowest temperatures. Finally, Figure 2-9 compares average monthly temperatures at Corpus Christi with those from the Middle Coast. Corpus is warmer during several months, though the differences are minimal.

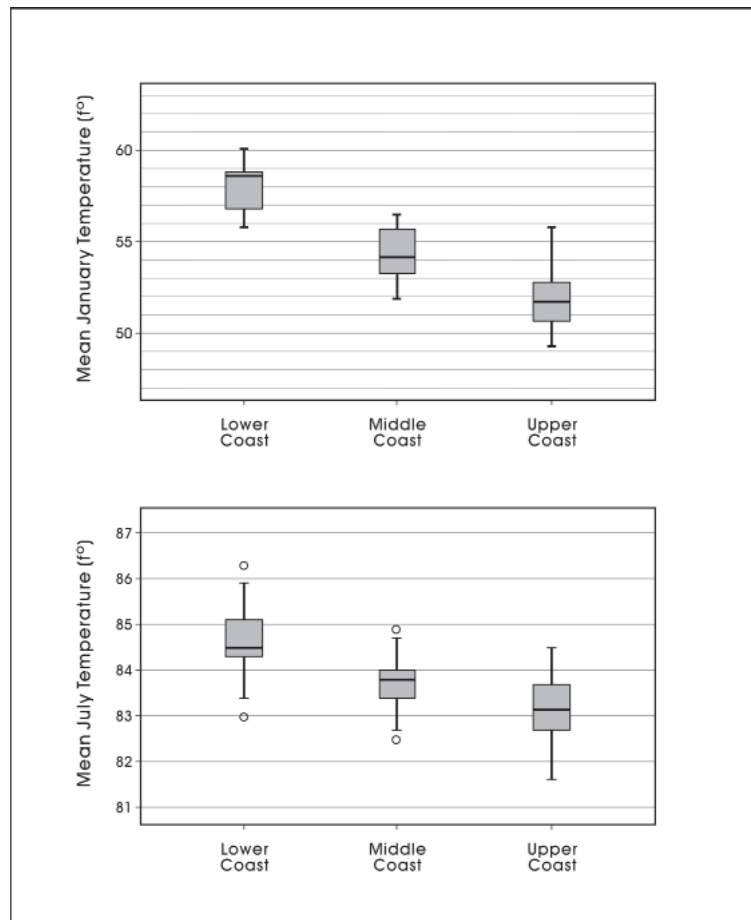


Figure 2-8. Mean January and July temperatures in the Texas Coastal regions. Based on data from NCDC (2004).

Overall, then, the modern climate of the coastal regions are wetter and warmer than the rest of the state. Along the coast as such, the Lower Coast is significantly warmer and drier relative to the Middle and Upper sections. The pattern at Corpus Christi and falling within the Middle Coastal region, is more reminiscent of the Lower Coastal pattern, a function of the southern location of this weather station within the Middle Coast. The climate reflected in the Corpus Christi data sets is also highly variable, especially with regards to precipitation. Mean annual evaporation rates range from 90-115 cm but may also reach 150 cm/year during drought periods. Because evaporation rates exceed precipitation, the region is under a constant moisture deficit. Summer high temperatures average 33.3°C, while winter lows average 8.3°C.

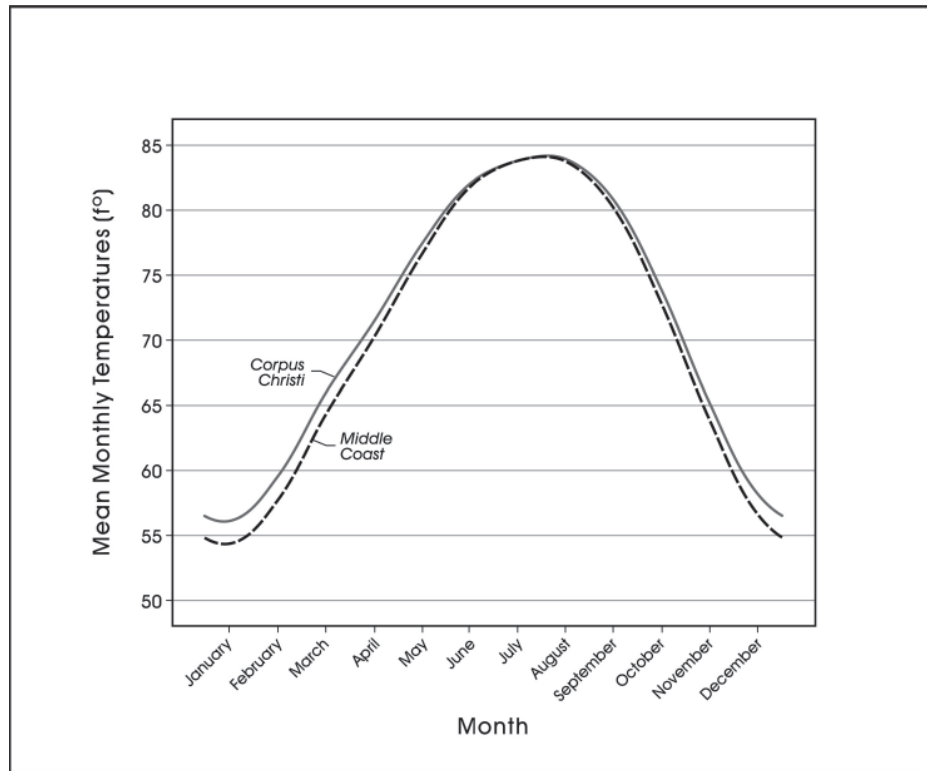


Figure 2-9. Relationship of mean monthly temperature at Corpus Christi and the average for the Middle Coast region. Based on data from NCDC (2004).

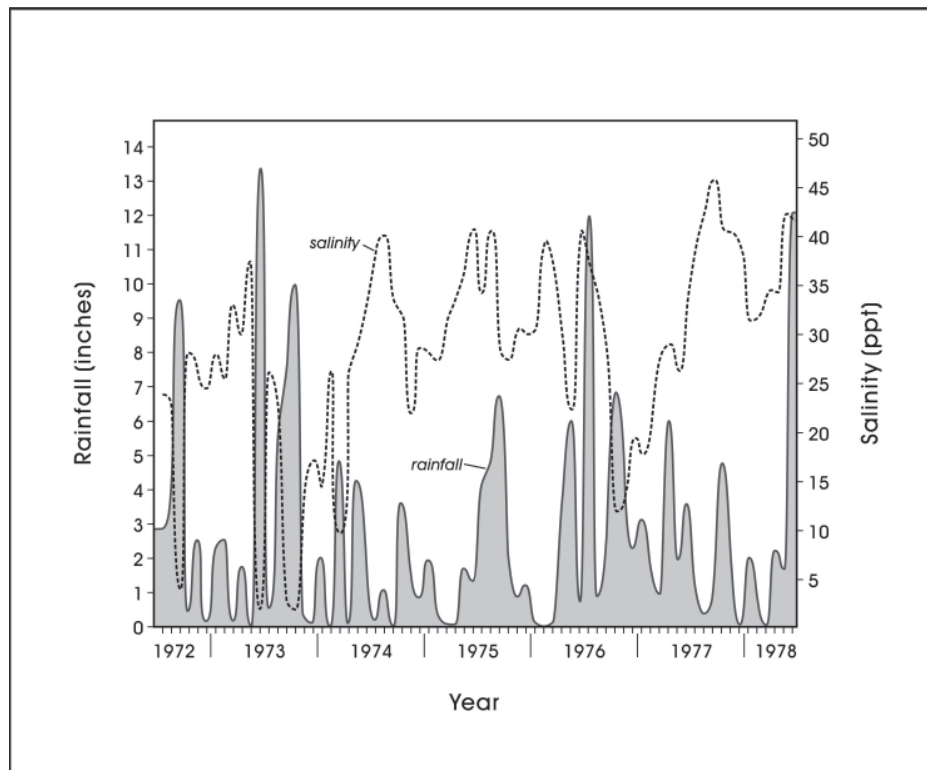


Figure 2-10. Seasonal and annual relationship between rainfall and salinity in Oso Bay during 1972-1978. (Based on data from Brown et al. 1976).

temperatures and low tides for several days following the passage of the front (Chabreck 1990). It is the combination of rainfall, temperature, and salinity that strongly influence habitat conditions for fish and molluscs.

Environmental Structure

Corpus Christi Bay is only one of several large bays along the Texas shores of the Gulf of Mexico. It lies within the Coastal Prairies sub-region of the Gulf Coastal Plains physiographic region of Texas (Brown et al. 1976). The Coastal Prairies sub-region forms an approximately 40-mile strip of land that parallels the shoreline, from Galveston Bay at the eastern edge of the state and progressing westward and south to the Rio Grande River (Figure 2-11).

In environmental and biological terms, the 40-mile-wide strip of the Coastal Prairies that is defined as the Corpus Christi area (Figure 2-11) by the Bureau of Economic Geology (Brown et al. 1976) encompasses all or parts of seven counties (Aransas, Bee, Jim Wells, Kleberg, Nueces, Refugio, and San Patricio). This area encompasses the northern portion of the Laguna Madre and Laguna Larga at its southern end and portions of Copano and Aransas Bays at its northern end. Mustang Island, St. Joseph Island and the northern end of Padre Island are the barrier islands off the mainland (Figure 2-11).

The BEG survey of the Corpus Christi area (Brown et al. 1976) identified a mosaic of 36 natural environmental zones and biological assemblages (Figure 2-12). These natural environments and biological assemblages can be grouped into habitats. A habitat is defined as an aggregation of biological communities that contain a relatively homogenous group of resources that are the product of long-term evolutionary dynamics and forces (i.e., the relationship between primary biomass and secondary biomass, primary productivity versus species diversity and the body size of the resident species), and as an aggregate are more similar to each other, based on the niche requirements of the resident biological assemblages and complementarity, than to adjacent habitats.

The 36 zones defined by Brown et al. (1976) can be divided into two broad categories: submerged or partially submerged environments (n= 15),

and terrestrial environments (n=21). Of the 15 submerged/partially submerged settings, 11 may have been accessible to humans. Ten of these submerged or partially submerged natural areas can be combined to form the highly diverse Bay-Estuary-Lagoon Habitat (Table 2-1). Fresh and brackish water ponds and lakes, the eleventh submerged setting, is found dispersed throughout the prairie uplands and it is therefore included within the terrestrial habitats mentioned below (Table 2-2). Table 2-1 summarizes those characteristics of the ten submerged environments that make up the Bay-Estuary-Lagoon Habitat that are most likely to impact the species composition and density of taxa present in these environments. In terms of area, the largest component of the Bay-Estuary-Lagoon Habitat is the open bays (165.9 mi²), followed by grass flats at shallow bay margins (53.4 mi²), enclosed bays with reefs (38.0 mi²), and bay and lagoon margins (31.7 mi²; Brown et al. 1976:Table 7). The Bay-Estuary-Lagoon Habitat covers a combined 385.7 mi.² of the Corpus Christi area. A variety of fish, snails, clams, oysters and molluscs and a multitude of bird species inhabit these environments.

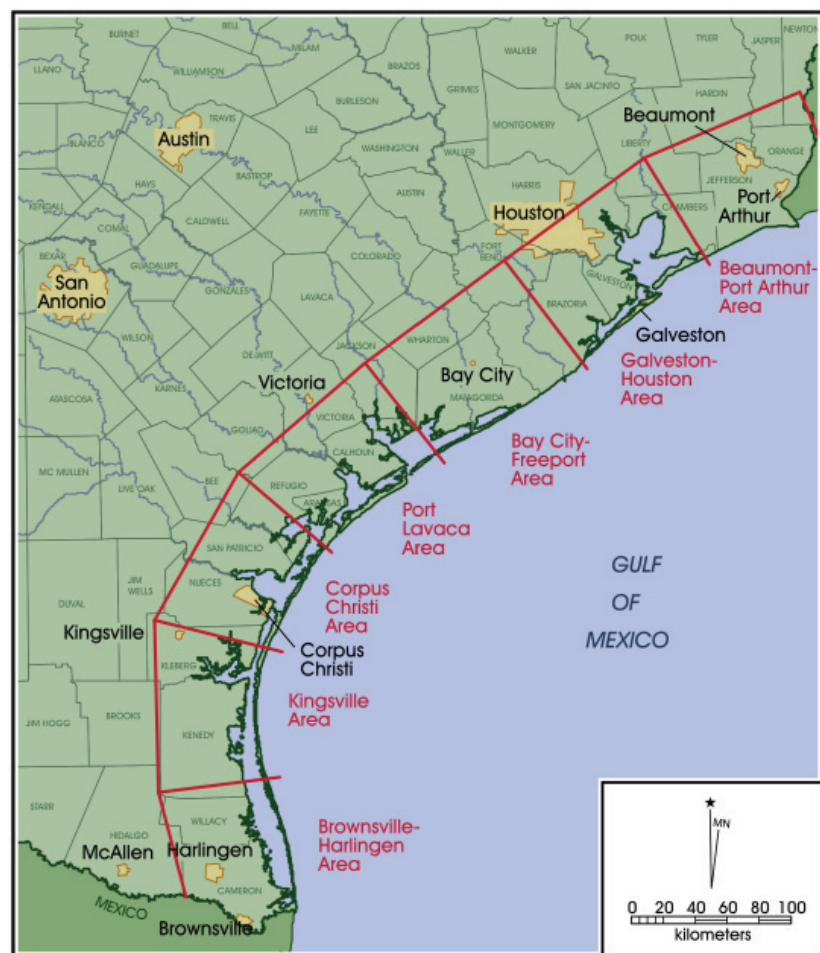


Figure 2-11. *The Coastal Prairies subregion of the Gulf Coastal Plains geographic regions, of the Gulf Coast of Texas (after Brown et al. 1976).*

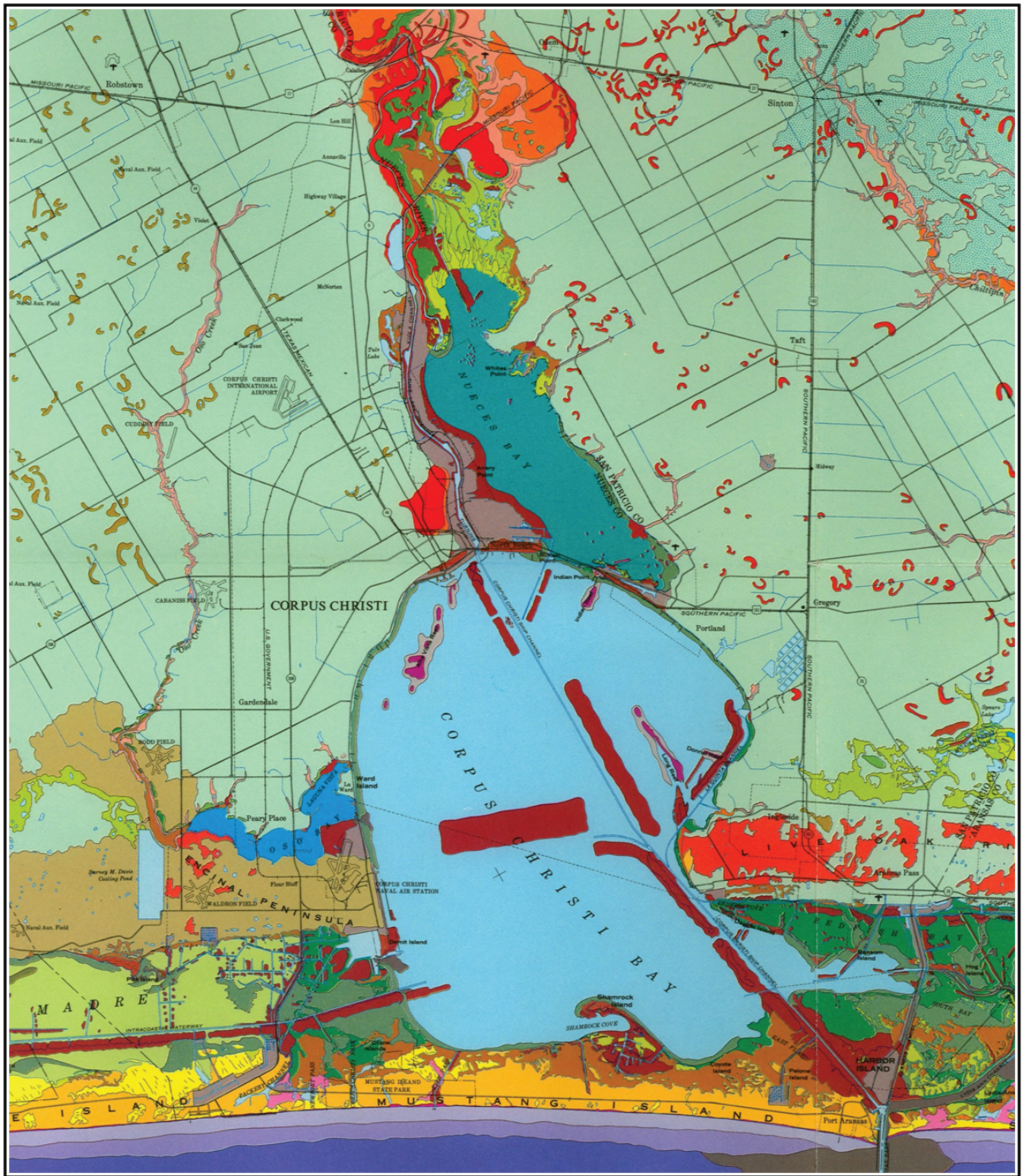


Figure 2-12a. Natural environmental zones and biological assemblages in the vicinity of Corpus Christi Bay, identified by the Bureau of Economic Geology (Brown et al. 1976).

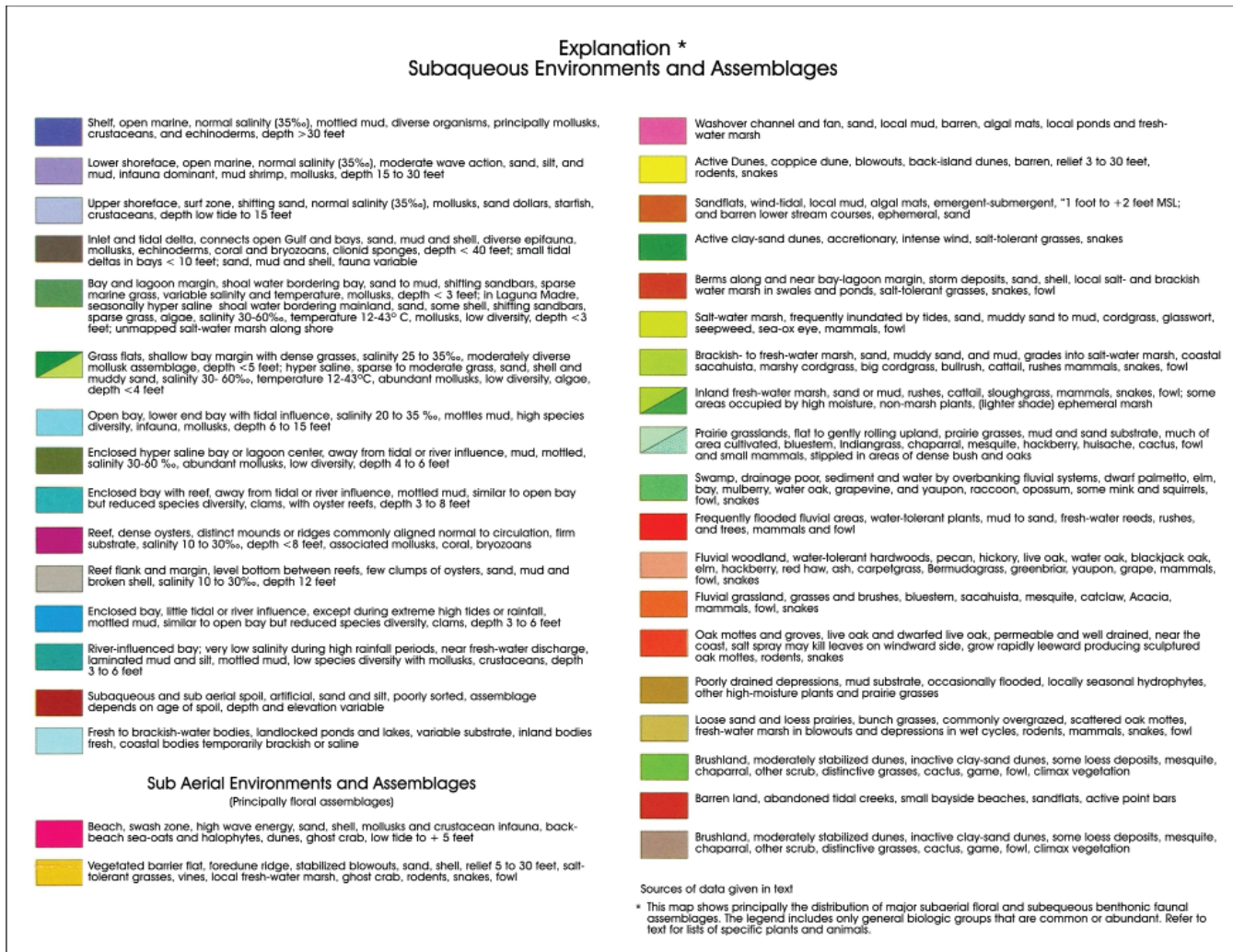


Figure 2-12b. Color key for Figure 2-12a.

Table 2-1. General Characteristics of the Submerged or Partially Submerged Natural Environments and Biological Assemblages

	Natural Environment/ Biological Assemblage	Principal Characteristics
Bay Estuary-Lagoon Habitat	Inlet and tidal deltas	Connect open Gulf and bays; sand, mud and shell substrate; clams and oysters; small tidal deltas < 10-feet deep;
	Bay and lagoon margins	Sand and mud substrate; clams; salinity ranges from 30-60 ppt; temperatures range from 12-43° C; < 3-feet deep; low diversity of mollusks;
	Grass flats	Shallow bay margins w. dense grasses; salinity ranges from 25-60 ppt; temperatures range from 12-43° C; < 5-feet deep; moderate to high mollusk diversity;
	Open bays	Deeper ends of bays under tidal influence; mottled mud substrate; salinity range from 20-35 ppt; 6-15-feet deep; high species diversity;
	Enclosed hypersaline bay and lagoon centers	Located away from tidal or river influences; mottled mud substrate; salinity range from 30-60 ppt; abundant mollusks but low species diversity; 4-6 feet deep;
	Enclosed bays with reefs	Located away from tidal or river influences; mottled mud substrate; clams and oyster reefs; low species diversity; 3-8 feet deep;
	Reefs	Firm substrate; salinity range from 10-30 ppt; dense oyster populations and associated mollusks; < 8-feet deep;

Table 2-1. Continued...

	Natural Environment/ Biological Assemblage	Principal Characteristics
Bay Estuary- Lagoon Habitat	Reef flanks and margins	Level bottoms between reefs; sand, mud and broken shell substrate; salinity range 10-30 ppt; clumps of oysters; depth <12-feet;
	Enclosed bays	Little tidal or river influence except during high tides and rainfall; mottled mud substrate; clams; low-moderate species diversity; 3-6 feet deep;
	River influenced bays	Near freshwater discharge zones; very low salinity during high rainfall periods; mud, mottled mud and sand substrate; low mollusk species diversity; 3-6 feet deep;

Table 2-2. General Characteristics of the Terrestrial Natural Environments and Biological Assemblages, Including Fresh and Brackish Water Ponds That Occur in All Habitats

Habitat	Natural Environment/ Biological Assemblage	Principal Characteristics
Beach and Beach-Proximate	Beaches	Exposed to high wave action and low tides, may cover up to 5-feet; mollusks and crustacean fauna;
	Vegetated barrier flats	Relief from 5-30 feet; salt-tolerant grasses, vines, chenopodium; mesquite, hackberry, prickly ash; rodents, snakes and waterfowl;
	Washover channels	Mostly barren; scattered stands of cordgrass, glasswort, saltwort; contain localized ponds and freshwater marshes;
	Active sand dunes	Largely barren; scattered stands of salt-tolerant grasses; relief 3-30 feet; rodents, snakes;
	Sand flats	Occasionally covered by water from wind and tidal action; may be only 2-feet amsl; cordgrass;
	Active clay dunes	Accretional in nature and subject to intense winds; salt-tolerant grasses (buffalograss, whitegrass) help stabilize surface; mesquite, pricklypear and Spanish dagger common; snakes, small mammals;
	Berms at bay-lagoon margins	Localized salt and brackish-water marshes and ponds; salt-tolerant grasses; mesquite, salt cedar; snakes and waterfowl;
Beach-Fluvial Interface	Salt water marshes	Frequently inundated by tides; cordgrass, glasswort, saltwort, saltgrass, black mangrove; crabs, snails, waterfowl, racoons, and small mammals;
	Brackish to freshwater marshes	Coastal sacahuista, marshy cordgrass, big cordgrass, bullrush, cattail, rushes, mammals, snakes, waterfowl;
	Inland freshwater marshes	Coastal sacahuista, seacoast bluestem, balsamgrass, buffalograss, Indiangrass, husache, mesquite, hackberry, oaks; mammals, snakes, waterfowl;
Prairie-Riverine	Swamps	Dwarf palmetto, elm, mulberry, water oak, yaupon, grape, racoon, opossum, some mink, squirrel, snakes, waterfowl;
	Frequently flooded fluvial areas	Water-tolerant plants; freshwater reeds, rushes, cattail, and trees (i.e., willow); various mammals similar to swamps; waterfowl;
	Fluvial woodlands	Water-tolerant hardwoods, pecan, hickory, live oak, water oak, blackjack oak, cedar elm, anacua, hackberry, red haw, ash; carpetgrass, Bermudagrass, greenbriar, yaupon, grape; squirrels, fox, racoon, opossum, other small mammals, rodents, fowl, snakes; March-June and August-October large numbers of birds;
	Fluvial grasslands	Bluestem, coastal sacahuista, mesquite, catclaw, rabbits, opossum, skunk, fox, quail, other fowl, small mammals, snakes;
Prairie	Oak mottes and groves	Live oak and dwarfed live oak; small rodents and snakes;
	Brushland	Mesquite, chaparral, pricklypear, Spanish dagger, buffalograss, whitegrass, coastal sacahuista, mesquitegrass, bermudagrass; deer, turkey, quail and other fowl;
	Prairie grasslands	Bluestem, Indiangrass, chaparral, mesquite, hackberry, husache, pricklypear, small mammals, waterfowl, localized dense areas of brush and oak mottes;
	Loose sand and loess prairies	Scattered live oak mottes ; blowouts that formed areas of freshwater marshes; bunch grasses; small rodents, mammals, snakes, fowl;
	Poorly drained depressions	Prairie grasses (bluestem, buffalograss), water-tolerant plants (sloughgrass, rush); distribution and size tied to seasonal wet periods;
Within all terrestrial habitats	Fresh and brackish water ponds and lakes	Landlocked pond and lakes; variable substrate; inland bodies fresh and coastal bodies brackish water;

Of the 21 terrestrial natural environments and biological assemblages, 19 would have been available to human inhabitants of the coast. These can be combined into four habitats, the Beach and Beach-Proximate Habitat, the Beach-Fluvial Habitat, the Prairie-Riverine Habitat, and the Prairie Habitat. Table 2-2 summarizes the general characteristics of these 19 terrestrial environments and biological assemblages that make up these four habitats as well as the fresh and brackish water ponds and lakes from the submerged category, because they are found within all four of the terrestrial habitats.

As can be seen in Figure 2-12a, areas near the coastline have a wide variety of inter-fingered natural environments and biological assemblages, such that any given area near the coast may be within one such environment but have several others closely adjacent. Fluvial-deltaic and marsh-swamp natural environment contains eight (42%) of the 19 natural terrestrial areas. Coastal beaches and barrier islands (barrier-strandplain and eolian natural systems) in combination contain seven (37%) of the 19 terrestrial assemblages present in the Corpus Christi area. Coastal prairie uplands tend to be more homogeneous containing only 3 of the 19 (16%) natural areas. They are dominated by grasslands, oak mottes and brush communities.

In terms of area, the largest terrestrial biological community in the Corpus Christi area is the prairie grasslands (1455.3 mi²), followed by sand and loess prairies (93.7 mi²), fluvial woodlands (61.1 mi²), frequently flooded fluvial areas (52.3 mi²), and fluvial grasslands (40.0 mi²). Saltwater marshes (10.7 mi²), brackish to fresh-water marshes (8.6 mi²), and inland fresh-water marches (29.9 mi²) also are a significant component in the Corpus Christi area (Brown et al. 1976). These 19 terrestrial environments, combined, are home to a rich mix of small and medium-bodied mammals, fowl, snakes, prairie grasses, oak mottes, and moisture-tolerant plant species.

As expected, the Prairie Habitat occupies the largest area (1,623 mi²), followed a distant second by the Prairie-Riverine habitat (154 mi²), the

Beach and Beach-Proximate Habitat (92 mi²), and the Beach-Fluvial Interface Habitat (49 mi²). Figure 2-13 is a graphic representation of the distribution of the five habitats we have defined in the vicinity of Corpus Christi Bay and that may also be present near other bays along the coast. In summary, within the coastal prairies geographic area, the Prairie Habitat (e.g., prairie uplands) constitutes the largest habitat although it has the least biological diversity. As one moves to low-lying parts of the coastal plains and beach-proximate settings, the diversity of biological assemblages increases but the extent of the habitats decreases. The Prairie-Riverine Habitat with its more diverse biological assemblages extends into the prairie grasslands providing linear corridors of greater biological diversity (Figure 2-13). Biologically, the most diverse communities are located at the interface of land and saline, brackish and freshwater bodies, and consist of the Beach-Fluvial Interface and Bay-Estuary-Lagoon habitats.



Figure 2-13. Distribution of five habitats defined for the Corpus Christi Area (after Brown et al. 1976).

Chapter 3: Cultural Chronology of the Texas Coast

Jennifer A. Thompson

Much of the known archeological record along the Texas Coast has been examined by Ricklis (1996, 2004; Ricklis and Blum 1997). He has addressed the cultural developments of coastal groups in an ecological setting fitting to the Human Behavioral Ecology framework used in this study. Therefore, we draw upon his work for much of the cultural history described in this chapter. The chronology begins with the Archaic Period rather than the Paleoindian Period because during the Paleoindian Period, the last glacial period was just ending, and ocean levels were so much lower that the coast line was still some distance from the modern coast line. Any Paleoindian sites recorded on the present-day coast would represent inland groups from a terrestrial/riverine adaptation rather than coastal groups, as what is now Corpus Christi Bay was probably not inundated until ca. 7500-7000 B.P., during the Early Archaic period (Brown et al. 1976:20-21).

Archaic Period

The Archaic Period began around 8,000 B.P. and lasted until the appearance of bow and arrow and ceramic technology between about 1500 and 1000 B.P. This long period has been divided by researchers into three sub-periods, the Early, Middle, and Late Archaic.

The Early Archaic

Ricklis (2004:161) cites two major periods of occupation during the Early Archaic. The gap between these two periods, roughly 6800 to 5900 B.P. is represented by few archeological sites in the area (Ricklis 2004:164). Ricklis believes that geologic evidence suggests that sea level may have been rising rapidly at this time (Paine 1991; Ricklis 2004:176; Thomas and Anderson 1994), so that the food-rich shoreline shallows would have been inundated, and new shallows had not yet developed as a result of silting.

Initial Early Archaic Occupation Period: 7500-6800 B.P.

The earliest known occupation period has been dated to 7500-6800 B.P. (Ricklis 2004:161:161-162) at a few sites overlooking Nueces Bay. Sites 41SP136, 41SP153, 41NU266, and 41NU281 (Ricklis 1993, 1995b) returned dates from shell to this initial Early Archaic occupation (Ricklis 2004:160, Table 5.2). Site 41NU266 is located on a bluff above the Nueces River delta, while the others are located on a bluff

above a cove of Nueces Bay. The large number of estuarine molluscs associated with these dates suggests strongly that the incursion of marine waters within the narrow valley had progress at least as far as 41NU266 and there had been enough time for shallows to develop at or near the mouth of the river.

These components are thin but dense with oyster shell within Holocene deposits, often sitting directly on the Beaumont Formation. The excavations at these sites consisted of testing, and the units containing the Early Archaic materials were widely scattered. The excavated strata contain few artifacts within the dense, though thin shell lenses, suggesting that they were short term camps visited by small groups of people, likely there to harvest the shellfish. The small number of artifacts cannot support much inference about the material culture. Debitage and a core were recovered from 41SP153 and utilized flakes were recovered from 41NU266. Edge-flaked sunray venus shell tools were identified (Ricklis 1993, 1995a). These tools apparently continued in use to the Late Prehistoric Period (Ricklis 2004). Reports from these sites suggest that either bone preservation is very poor, or that the inhabitants of the camps did not expend much effort in hunting and fishing.

Terminal Early Archaic Occupation Period: 5800-4200 B.P.

This Terminal Early Archaic period is better documented than the earlier occupations and well dated. Twenty-four radiocarbon assays date fifteen discrete stratigraphic components from the Coastal Bend on the shores of Nueces Bay, Baffin Bay, the Nueces River delta, The Lavaca River delta, and Ingleside Cove (Ricklis 2004:Tables 5:1, 5:2). As in the Initial Early Archaic occupation, most components of the Terminal Early Archaic are thin shell lenses, with few bones and other artifacts (Ricklis 2004). More components are recorded for the Terminal Early Archaic period occupation, though sites from both this and the earlier occupation rest on the eroded surface of the Beaumont Formation and could have eroded during the Middle Holocene (Ricklis 2004:165). The McKenzie Site, 41NU221, located about 3 km upstream of 41NU266, contains a large thick shell deposit with hundreds of flakes, five dart points, a stone end scraper and shell tool (Ricklis 1988). At the Means site, 41NU184, extensive excavations produced debitage and some surface dart points near a shell midden (Ricklis and Gunter 1986). The site also contained a series of post molds in a semi-circular configuration, interpreted by Ricklis and Gunter (1986) as

the remains of a structure. Excavations at site 41SP156 on the Nueces Bay revealed an occupation dating from 5592 to 4614 B.P. (Ricklis 1993). An Andice barb and a possible re-used Gower point were found within the dated shell stratum. Site 41NU267, also on Nueces Bay, produced deep deposits containing early triangular points with oyster shells dating to 4986-4862 B.P. (Ricklis 1995b).

Otolith remains at many sites dating from the end of the Early Archaic indicate the presence of estuary fishes. Some identified species include black drum, spotted sea trout, redfish, Atlantic croaker, and catfish. These, along with the presence of dart points and shell lenses, indicate that some hunting likely contributed to a diet that also included fish and shellfish. Shell studies from oyster lenses on the Nueces Bay shore suggest shellfish gathering occurred during the winter and spring months, when plants were scarce and game was lean. Results of otolith analysis from 41NU221 (Ricklis 2004:165) indicated similar dependence on fishing at this site.

The postulated highstand with sea level about a meter higher than it is today was taking place at this time (Paine 1991; Ricklis 1997:7). This may explain why 41NU221, located about 6 km from the current mouth of the river, contained oysters (*Crassostrea virginica*). Oysters require water levels at least 8 feet deep and a minimum salinity of 15-40%. In addition, there were large numbers of the Brown Rangia (*Rangia flexosa*), a clam with a normal salinity tolerance that ranges from 5-19%.

Middle Archaic

Ricklis (2004: 164-165) defines another gap in the archeological record on the Central Texas coast that occurs during the Middle Archaic between 4200 and 3100 B.P. He believes that shallow water and marshes may have been reduced affecting the abundance of available aquatic resources to coastal groups. The dearth of Middle Archaic archeological sites on the coast is not likely due to sampling. Many sites have dated zones of occupation above and below the Middle Archaic geologic deposits (Ricklis 2004). This component was seen at 41NU2 with dated features falling between 4200 and 3500 B.P.

Late Archaic

By the beginning of the Late Archaic, the current coastal environment of bays and lagoons were beginning to take shape. At about this same time, the postulated highstand would have ended (Paine 1991; Ricklis 1997:7). Sea levels stabilized, promoting formation of the barrier islands, which coalesced into islands similar to the modern forms sometime between

about 3000 B. P. (Ricklis 1995a:142) and about 2500-2000 B.P. (Morton and Paine 1984:12). The barrier islands allowed back-water shallows to develop landward, as river sediments filled the bays forming marshes and lagoons along the shoreline, around Nueces Bay, the Nueces River delta, and Oso Bay. The broad shallow waters were optimal for fish spawning and likely served as a catalyst for a booming fish population.

The availability of resources from these wetlands attracted coastal groups, as evident by the archeological record. Sites from the Late Archaic are larger, thicker, and include a greater quantity and variety of artifacts. Faunal bone assemblages consisting of more than 100 specimens occur for the first time and suggest that fishing and hunting were part of the Late Archaic subsistence strategies. An increase in the number of fish otoliths coincides with the formation of the barrier islands and the estuaries and marshlands that formed as a result (Ricklis 2004:165). The abundance of estuarine resources encouraged population growth among the human population, as well, and may explain the increased site density across the region and richness of the archeological deposits.

Excavated coastal sites with Late Archaic components include the Kent Crane site (41AS3), Mustang Lake (41CL3), 41SP43, and 41SP120 (Ricklis 1996:36-56). Bone, stone, and shell tools have all been recovered from Late Archaic sites as well as indirect evidence of basketry i.e., basketry-impressed asphaltum at the Holmes Site, 41SP120; (Ricklis 1996:41). Late Archaic components also contain the first recognized cemeteries along the Texas coast. The rise of cemeteries along with increased population may indicate increased territoriality associated with prime hunting or fishing grounds (Ricklis 2004:169).

Late Prehistoric Period

The Late Prehistoric period is marked on the Texas coast as it is throughout the state, by the arrival of the bow and arrow and ceramics (Ricklis 2004:171). Coastal groups continued to rely on the estuarine resources for subsistence at least part of the year, as evident in the amount of fish bone and mollusc shell present in archeological components. Identified fish species include black drum (*Pogonias cromis*), red drum (*Sciaenops ocellatus*), spotted seatrout (*Cynoscion nebulosus*), and Atlantic croaker (*Micropogonias undulates*), all of which are fish that spawn in or near Corpus Christi Bay, where young fish can find places to hide in the sea grass beds (Ricklis 1995, 1996). Shellfish, such as oyster (*Crassostrea virginica*), lightning whelk (*Busycon sinistrum*) and bay scallop (*Argopectin irradians*), continued to be present in Late Prehistoric shell lenses. Hunting is evident through the presence of arrow points and vertebrate faunal remains of deer and bison (Ricklis 1996).

Initial Late Prehistoric

Ricklis identifies two sub-periods of occupation within the Late Prehistoric. The Initial Late Prehistoric lasted from A.D. 1000-1300. The distinction is based on well-dated terminal Archaic stratum at A. D. 1000 beneath a discrete component with Fresno arrow points and abundant fish bone dated from two shells to A.D. 1022-1212 and A.D. 1024-1209 (Ricklis 1993). Initial Late Prehistoric Scallorn arrow points have been recovered from the Mustang Lake site, the Kendrick's Hill site, and at the Anaqua site (41JK8), where they were found with sand-tempered pottery (Ricklis 2004:171).

Terminal Late Prehistoric

The Terminal Late Prehistoric sub-period along the coast is defined by the emergence of a particular artifact assemblage, the Rockport Phase (A.D. 1300-1700), similar in many ways to the Toyah Phase assemblage of Central Texas. Both include Perdiz arrow points, small unifacial end scrapers, thin bifacial knives, chert drills, sandstone metates, and prismatic blades and cores. The ceramic assemblage sets the Rockport Phase apart from the inland Toyah assemblage of the Terminal Late Prehistoric. Rockport phase shell tools include conch and whelk adzes, awls, hammers and the edge-flaked Sunray Venus clamshell knives and scrapers that were seen in components dating 6000 years earlier. Shell beads and perforated shells have also been discovered in Terminal Late Prehistoric assemblages (Ricklis 2004).

The Rockport Phase ceramic collection contains few complete specimens though some inferences from sherds indicate that vessel forms include bowls, jars, and ollas sometimes decorated with asphaltum, incised, and crenulated along the vessel rim. These Rockport Phase ceramics have been found along the coast in association with bone tools, bone beads, and shell tools from Matagorda Bay and the lower Colorado River to northern Baffin Bay and up to 40 km inland with Toyah Phase artifacts (Ricklis 2004:172). Suhm and Jelks (1962) defined three pottery types with the Rockport Phase ceramic tradition: Rockport Plain with no decoration; Rockport Black and Gray, with asphaltum bands painted both inside and out after firing; and Rockport Incised, decorated with geometric lines incised around the sub-lip area externally.

Ricklis (1990, 1996) revised the typology adding a fourth type and two varieties to the Suhm and Jelks (1962) typology. The Rockport Crenellated type has a plain body with a notched rim decoration and occasionally an asphaltum band at the lip. Rockport Black on Gray was divided into two varieties. Variety I includes wide mouthed bowls and jars with an external asphaltum band and no internal decoration. Variety

II includes small-mouthed neckless or constricted ollas, with asphaltum lines from the lip to the base and asphaltum coating the interior of the vessel (Ricklis 1990, 1996).

Subsistence seems to have shifted during the Terminal Late Prehistoric to more reliance on fish than shellfish and the inclusion of bison. Hunting camps with bison and deer faunal remains attributed to the Rockport Phase have been found inland overlooking drainages in prairie settings at 41NU221, 41NU193, 41NU255, 41NU46, 41SP159, 41SP160, 41SP161, 41SP167, and 41SP170. These sites are smaller with fewer artifacts than Rockport Phase coastal sites, which cover a larger area and have a relatively higher artifact density. Settlement patterns may reflect seasonal use of aquatic resources during fall and early spring at coastal camps and hunted prey from late spring through summer at stream valley hunting camps in the prairie. The larger coastal sites versus the smaller riverine prairie sites might also indicate population aggregation at the coast before dispersal to the prairie for hunting forays (Ricklis 2004).

Historic Period

Early European encounters along the Texas coast have provided us an ethnohistoric record of the indigenous groups, collectively known as the Karankawa, who lived roughly from Galveston to Corpus Christi Bay. The first extant record of European contact with natives on the Texas coast is from the Narváez expedition in 1528. An excellent recent short review of the story told by the treasurer of the expedition, Álvaro Núñez Cabeza de Vaca, and his three companions who arrived in Vera Cruz more than seven years after being shipwrecked on what was probably Galveston Island, is available at the *Texas Beyond History* website (Thoms 2008). The 1555 edition of *La relacion y comentarios del governador Alvar nuñez cabeça de vaca, de lo acaescido en las dos jornadas que hizo a las Indias* is available, with facsimiles of each page and an English translation (Cabeza de Vaca 1555). *La Relación*, describes the Narváez expedition, the shipwreck off what was almost certainly Galveston Island, and the lifeways of the Native Americans who kept them alive during their first winter. It then describes the attempt by survivors of that first winter to find the way to Spanish-controlled Mexico. Although the route of their wanderings is still in dispute, it is possible that the first part of their trip was from the mainland near Galveston to an area that seems to have been the lower Guadalupe River and San Antonio Bay. After this, they turned inland. Of the eighty men who came ashore on Galveston Island, only four had survived, eventually reaching other Spanish speaking people (Thoms 2008).

Another account of Karankawa lifeways, more than a century later, comes from the interrogation of two French brothers,

Pierre and Jean-Baptiste Talón, who along with several siblings, were children saved by Karankawa women during the massacre of the last remnants of the French colony Fort St. Louis, near Matagorda Bay (Bell 1987; R.B. Campbell 2003:41-46). An account of the colony by LaSalle's friend Henri Joutel (Foster 1998) is a more complete story but is less interesting from an ethnographic perspective, as the story told by the brother years later in France relates the lifeways of both the Karankawa and the Cenis, an inland farming Native American group. The brothers were taken away from their lives with the Karankawa and Cenis during the expedition by the de Leon expedition, sent to find the French encampment after learning of the La Salle expedition (R.B. Campbell 2003: 46). The Spanish, who had come to challenge what they considered a French land grab, returned the boys to Mexico, where they were held until adulthood, at which time the two oldest boys joined the Spanish Navy, and eventually made their way back to France (Bell 1987).

Another French account comes from François Simars deBellisle, who was marooned near Galveston Bay in 1719 (Weddle and Lemée 2003). DeBellisle was a French officer with the French West Indies Company sailing for Louisiana. Missing the Mississippi River, the ship landed in Galveston Bay where Bellisle and four others went ashore to assess their location and were left. Only deBellisle survived until he was encountered by an Atákapán band, who spoke one of the Karankawa dialects. He remained throughout the summer and the winter of 1720. A letter passed from tribe to tribe reached the French colony at Fort Saint-Jean-Baptiste, who sent a few Hasinai, and inland group, to rescue deBellisle. DeBellisle then remained with the Hasinai for several months before finally reaching the French settlement at Natchitoches in February 1721 (Weddle and Lemée 2003:23).

Other accounts of the native Texas population come from the Spanish reports from missions. In 1690, as a result finding what was left of the LaSalle colony, the Spanish began securing the northern border of New Spain, expanding their interests in east Texas to counter any French expansion across the Mississippi River. The Marqués de San Miguel de Aguayo scouted Texas for possible mission locations for coastal tribes (Ricklis 1996). The relationship between the Spanish colonists and coastal Native groups was hostile, likely fueled by real and perceived threats on both sides to life, land and property. As they had done, with limited success, with the

farming communities in East Texas (Ricklis 1996:126), the Spanish attempted to use missions as a method to convert the Native Americans to Catholicism, and, hopefully, to Spanish dominance. The first near-coast Spanish mission, Nuestra Señora del Espíritu Santo, begun in 1722, originally occupied near the old Fort Saint Louis, which had been rebuilt by the Spanish. This mission later relocated with the presidio to the Guadalupe River valley near what is now Victoria, in 1726, and then to Goliad in 1749. Nuestra Señora del Rosario was established in the mid-1700s, again to convert the coastal indigenous population near La Bahía and to exert Spanish control over the coast. Again, the mission met with little success in converting the Karankawa groups who abandoned the mission by the 1780s (Ricklis 1996:150). Nuestra Señora del Refugio mission was established in 1795 as a final attempt at assimilation of the native population and met with some success with conversions to Christianity. During this time, the Karankawa agree to a treaty with the Spanish that was the start of 40 years of peace (Ricklis 1996:152-153).

European colonization of the eighteenth century changed the demographics of the indigenous populations everywhere across the continent. In coastal Texas, Karankawa populations began to decline, largely due to the introduction of new diseases, and began to merge with interior populations displaced to the coast by southward expanding Apaches and Comanches. Ricklis (1996:139; Figure 26) identifies a period of several decades after 1685 that saw a seventy percent reduction in Karankawa population followed by a period of relative stability until the early nineteenth century when Anglo-Americans began receiving land grants on the coastal prairies near traditional Karankawa territory. The Karankawa population declined again until around 1840 when numbers dwindled to less than 100 (Ricklis 1996:140; Table10). Remaining survivors then likely moved south into Mexico and merged with other groups (Newcomb 1983).

The French and Spanish accounts of indigenous populations prior to and during colonization provide information on group size, camp locations, and subsistence strategies that support the findings in the archeological record of pre-contact sites. Both suggest a pattern of large groups camping on the shores during winter and spring, then dispersal inland in smaller groups to the prairie for hunting during summer and fall. These subsistence and land use strategies are summarized in greater detail in Chapter 7.

Chapter 4: Field and Laboratory Methods

Barbara M. Meissner

The methods for recording archeological deposits exposed during construction were determined by TxDOT. TxDOT was primarily concerned with avoiding or carefully monitoring impacts to any human interments. The methods for excavating any burial located were determined at the beginning of the project and did not change. The recording of artifacts and features seen during monitoring in the construction trenches changed during the course of the project, and the changes will be detailed below.

Provenience Control

For a number of reasons, providing for provenience control during this project was very difficult. In addition to the overall size and topographic layout of the site, the fact that heavy equipment would be rolling over virtually every part of the ROW within the site boundary made setting a single datum impossible. From the beginning, in 2000, Cindy Tennis, who was the Project Archeologist at the time, designated the center line of the new road as the horizontal datum. This line, carefully surveyed by the construction subcontractor, Haas-Anderson Inc., was marked every 20 meters either by paint on the paved areas or by labeled stakes on open ground. It extended from one end of the project, at the intersection with the access road on the north side of State Highway 358, to the center of the intersection with Ocean Drive, following Ennis Joslin roughly north and south. Another centerline ran the extent of the project on Ocean Drive running roughly east and west. The centerline was frequently re-surveyed, whenever the wear-and-tear of construction activities made the pavement marking illegible or the stakes were damaged or removed by vandals. The road centerline provided us with secure horizontal provenience that would remain the same throughout the project, and could be reconstructed in the future by knowing that the inner edge of the curb on the new road was 11 m from the centerline. Any given point along the centerline could easily be identified by measuring from the nearest known identified point. The provenience of any given location could thus be described as a certain distance, in a given cardinal direction, from the centerline at a specific point along that line. Thus a point might be described as being 11.63 meters east of the centerline at station 33+579.51 on the Ennis Joslin centerline, or 23.2 m north at station 3+425 on the Ocean Dr. centerline. While this method ideally had the potential to measure accurately to the nearest centimeter, in practice, getting the exact location from the centerline could sometime be difficult and, not infrequently, dangerous, due to the extremely heavy traffic. For ordinary purposes, then, distance from the centerline was measured from the edge of the trench. Only for very important

findings, such as the six human burials, was the distance from the centerline measured directly.

In the field, horizontal provenience was determined by depth below ground (bgs) or road surface (brs), with notes made whether this was the surface of the old road or the new road. The ground and road surfaces were mapped both before and after the new road surface was completed, so that an estimate of the elevation of any given point on the surface could be made.

Monitoring Trench Excavations

Throughout the project, TxDOT was meticulous in notifying CAR when construction activities that would break the ground surface were planned. TxDOT officials made it very clear to all subcontractors working at the site that they were to follow instructions from the monitors, when, for instance, monitors requested that an activity be stopped so that the monitor could take a closer look at the walls or bottom of a trench or record a feature.

In order to allow time to deal with any human burials or important cultural materials that might have been found during construction, most of the utility trenches were excavated well before, sometimes months or even years before, the plans called for actual placement of utilities in the trench. Trenches were, for the most part, dug with a Gradall, at a rate of roughly 2-3 cm at a time, in sections roughly 3.5 m long. Most of the long trenches that ran parallel to the road were either 1 or 3 meters wide. A backhoe was used when the trench was to be less than 1 meter wide, or when the compacted road base was particularly dense and difficult to break up. A large track hoe with a 2.4 m bucket was used to re-dig trenches that had already been cleared. For safety reasons, once trenches were cleared they were backfilled, either at the end of the day, or when a section was completed, depending on the location of the trench.

In general, the trenches were dug to the depth of the utility grade plus about 10 cm or to the top of the Pleistocene Beaumont Clay Formation (see Chapter 2), whichever came first. The TxDOT superintendent and/or the foreman for the subcontractor provided both the location of the trenches and the depth of the utility grade at various locations. The depth of the trenches varied from about 50 cm to as much as 3 m. When trenches extended below the OSHA safety standards,

CAR personnel did not enter unless the walls of the trench were stepped down on each side until the lowest part of the trench was shallow enough to meet the safety standards (Figure 4-1). In some cases profiles of the lower reaches of a deep trench were drawn by standing in the partially excavated next segment of the trench when it was no deeper than OSHA safety standards and measuring deeper soil changes from that point.

CAR monitors were always present when the ground surface was penetrated. Monitors stood as close as was safe to the edge of the trench in a location that allowed as much of a view of the bottom of the trench as possible (Figure 4-2). The long trenches were not always dug in a continuous way, particularly on the east side of Ennis Joslin. Due to various exigencies inherent in the attempt to completely rebuild a road while it was still being heavily used, and without blocking the many residents from their driveways, the trenches were often dug in sections of various lengths. Beginning in May 2003, a “trench map” was created as the longer trenches were cleared, so that it was possible to keep track of what portions



Figure 4-1. Trench with sides stepped down to allow CAR monitors to enter safely.

of each trench had been completed. Features were drawn to scale on the map, and location of buried utilities or isolated finds were noted. The trench maps also became invaluable as Trench I was dug next to Trench B, allowing correlation of partial features and Occupation Zones (OZs) across the trenches (see Chapter 6).

During Phase III (see below), when trenches were dug and utilities placed immediately, and when long stretches of road bed were being graded, special attention was given to avoiding anything that would slow the progress of the construction. For instance, profiles, when needed, were sometimes drawn while the subcontractor’s workers were away at lunch.

Recording Features

Phase I: August 15 to October 3, 2000

The original methods requested by TxDOT for monitoring the trench excavations, recording features, and the excavation and collection of any human remains encountered were discussed in mid-August 2000, the day before the excavations of the first trenches were to begin. At that time, the intention was to clear the entire ROW not yet paved (C. Tennis, Field Notes, 8/14/00), leaving only the sediments under the road itself to clear later. Nancy Kenmotsu, the representative of the Environmental Affairs Division of TxDOT, in meeting with the Project Archeologist and Dr. Robert Hard, at that time Director of CAR and Principal Investigator, stated the following requirements for CAR staff working at the site:



Figure 4-2. Monitoring trench excavation near the location where it is believed Jackson’s 1933 excavations took place. Facing east.

1. All TxDOT safety regulations were to be followed. All federal Occupation Safety and Health Administration (OSHA) regulations were to be followed.

2. As the project had the potential to encounter human burials, and possibly many human burials, and given the sensitivity of the situation if that were to occur, it was imperative that dissemination of information to the public and press flow through TxDOT. All questions from the news media and the general public regarding human burials were to be passed to the person at the Corpus Christi TxDOT District responsible for public relations.

3. The protocol in the event that a human burial was encountered was as follows:

a. The Project Archeologist was to call the TxDOT superintendent of the project and inform him of the situation. All work in the immediate area of the burial was to cease immediately and the area should be closed off with barricades, if it was deemed necessary. In any case, an attempt should be made to keep the sight of the burial from all but authorized personnel.

b. The Project Archeologist, and/or physical anthropologist on the site should clear enough of the bone to determine if it was human, or was likely to be human, and if so, to estimate how likely it was that the burial was intact as well as the position of the still-buried portions of the remains.

c. The Project Archeologist would then call the head of the TxDOT District Office in Corpus Christi, and the Director of CAR with this information. A discussion of the particular situation would then be initiated between the Director of CAR and TxDOT-ENV offices in Austin.

d. Once a protocol for dealing with the specific situation had been worked out, the Project Archeologist would make whatever arrangements were necessary to complete the excavation of the burial, being sure to keep the TxDOT District Office and the TxDOT-ENV office informed of progress. If it was necessary to leave the burial overnight, arrangements were to be made to protect it to insure that it could not be vandalized (i.e., hiring off-duty police as guards).

e. The human remains were to be excavated as quickly as possible while maintaining professional archeological standards of data retrieval. When removed, the bones were to be protected from any possible damage and returned to the CAR laboratory as soon as possible.

4. The protocol for recording non-burial features was as follows:

a. The feature was to be exposed, photographed, drawn, and mapped. A sample of the fill was to be excavated

in a 1 x 1 m unit. All material from this excavation was to be screened through ¼" (0.64 cm) hardware cloth and all cultural materials, including vertebrate and invertebrate fauna, was to be collected, and returned to the CAR lab for processing and analysis. If the feature was very large, more than one excavation unit would be opened.

b. A bulk soil sample would be taken of each feature.

5. The trench wall would be profiled every 5 m and an attempt made to tie features to the stratigraphy.

Excavations began on August 15, 2000 on the west side of Ennis Joslin at station 33+680. A few days later, the project engineers informed TxDOT that the plan to excavate and backfill everything within the ROW would leave the walls of the undisturbed sediments beneath the pavement unsupported, and would make the road very susceptible to collapse. TxDOT informed CAR that only the actual utility trenches would be excavated from this point forward.

Shortly after the work began, in an attempt to allow a faster pace of excavation, it was decided that only concentrations of shell or animal bone that included charcoal were to be defined as features. The following designations and treatment protocols were to be used during the monitoring:

1. A **feature** was defined as a dense concentration of shell and/or animal bone that contained charcoal. Features identified were to be hand-excavated in 10 cm levels, screened through ¼" (0.64 cm) mesh, and soil samples were to be collected for floatation. Feature forms were to be completed detailing provenience, contents, context, and size. A profile of the trench wall adjacent to the feature was to be drawn. Photographs were taken, and drawings made when appropriate.

2. A **concentration** was defined as a deposit of shell and/or animal bone that was less dense than a feature and did not contain charcoal. Concentrations were to be hand-excavated, but not screened. Provenience information and a brief description of the concentration and its context were to be recorded in the field notes and collection-bag log. Soil samples were to be collected for humate dating and/or floatation. Photographs were to be taken, and drawings made when appropriate.

3. A **scatter** was a collection of shell and/or animal bone that was not in a concentration or was reasonably diffuse. Provenience information and a brief description of the scatter and its contents were to be recorded in the field notes and collection-bag log. Samples of the scattered shell or bone and soil samples for floatation were to be collected if considered appropriate.

At the end of the period, in early October, a total of 16 features had been excavated and notes and soil samples taken from a few “concentrations” and “scatters” but these were a very small part of the total volume of material returned to the CAR Lab.

On October 3, 2000, after a large water main that did not appear on city utility maps was broken during excavation near the southern end of the site, TxDOT decided to halt archeological monitoring until plans could be made to deal with the fact that utility maps were not accurate. In the mean time, TxDOT planned to concentrate on the southern part of Spur 3, near Highway 358, far outside the site limits.

Phase II: October 15, 2001 to April 23, 2003

After a year, CAR was once more called to 41NU2 to monitor ground disturbing construction at the site. Traffic was to be diverted to a temporary bypass built to the west of Ennis Joslin and the west lane, so that the east lane could be closed for construction. Later, the west lane was closed and traffic moved onto the two new lanes on the east side of Ennis Joslin while construction of the two west lanes was completed. During this phase, the protocol for recording non-burial cultural materials again was changed. Since almost all the sediments to be excavated at this time were believed to be previously disturbed, consisting largely of fill over old utility trenches, only notes were to be taken if cultural materials were observed in undisturbed sediments, as these areas would be disturbed again in a later phase and could be examined closely at that time.

During this roughly 18 month period, work was intermittent, running two or three weeks at a time to as little as a few hours with long gaps between. CAR monitored the following activities:

Utility Location. A large number of small trenches or pits, were dug to locate and map the location of utilities, especially lines making connection to the many houses and commercial buildings lining the roadway. These excavations almost exclusively dug into previously disturbed sediments.

Utility Relocation. In several locations, especially along the west side of Ennis Joslin south of the intersection with Alameda Street, existing utilities had to be temporarily relocated to the west so that a temporary one-lane bypass could be constructed. Most of this activity was at the center of the site, where the Beaumont clay was very near or at the surface, however, monitoring was still performed.

Road Grading. During this phase only the area intended for the bypass lane mentioned above was graded.

During the grading, a monitor walked behind the maintainer, closely observing the newly scraped ground surface. Later in the project, large areas of previously undisturbed sediments were graded. However, during this phase, the grading was exclusively over either previously disturbed sediments or undisturbed Beaumont clay.

Traffic Island Removal. The complex intersection of Ennis Joslin and Alameda Street was controlled by a series of traffic islands. To create the bypass, these islands had to be removed. Monitors were present during this process, in case previously undisturbed sediments would be impacted.

Clearing Trenches for Future Utility Placement. Although most of the monitoring during Phase II consisted of the activities listed above, four trench laterals that crossed Ennis Joslin between Alameda Street and the southern end of the site, were monitored in advance of utility placement with all four completed on one Saturday, when traffic disruption was minimal.

Phase III: April 24, 2003-November 3, 2006

Shortly before Phase III began, in a meeting with TXDOT officials, CAR staff were told that the protocol for recording features and cultural deposits was to be changed. The changes to the protocol, which did not apply to human burials, were as follows:

1. Features would no longer be excavated, nor would any sediments be screened in the field except during investigation of burials.
2. Features should be recorded as quickly as possible. The surface of a feature was to be cleaned enough to estimate size and composition. Scale drawings of the feature would be done only if it was very unusual, and should only be done in a fashion that would not delay the progress of construction. Photographs could be taken and should be the major form of documentation of the feature. Feature forms were to be filled out using estimates when exact measurements could not be determined. Soil samples could be taken if deemed necessary.
3. With the exception of temporally diagnostic tools, collection of artifacts and faunal remains should be minimal and should not impede the progress of construction.
4. Any bone observed during excavation, whether in a feature or as an isolated find, was to be carefully excavated by hand and examined to determine if it could be human.
5. Profiles of trench walls would no longer be drawn at 5 meter intervals. The Project Archeologist should determine when a profile was needed, but the number of profiles should be kept to a minimum, and be

accomplished as quickly as possible, making every effort not to slow construction.

Monitoring activities during Phase III included:

Trench Monitoring: During this phase, twenty trenches (including a re-alignment of the southern end of Trench B) were dug, and most were then re-dug during utility placement. Though most trenches had been previously cleared, the re-excavation and utility placement was monitored because of the possibility that either re-excavation or the extension of a trench to allow building of a manhole box would encounter a burial in previously undisturbed sediments.

Utility Removal. Once the new utilities were in place, old utilities were removed. This required excavation of a narrow trench with a backhoe to the level of the old utility lines. After the old lines were removed, the monitor occasionally observed areas of undisturbed sediments. These areas were profiled before the trenches were backfilled. In two instances, during the re-excavation of old utility trenches the monitor observed disturbed human remains (e. g. Burials 3 and 5) in the trench fill. As soon as this occurred, the protocol regarding human burials was implemented both in terms of notification procedures and excavation methods.

Monitoring New Road Grading. Old roadbed was removed and sediments were removed to the new sub-grade. The depth of the new grade cut into previously undisturbed sediments from 0 cm to approximately 35 cm. CAR archeologists monitored this grading by following the maintainer and noting presence of cultural debris. All observed bone was excavated by hand. When identifiable features or occupation zones were located, the work was stopped as briefly as possible so that notes and measurements could be taken. Otherwise, grading was allowed to proceed under careful observation. Once the new grading had been completed, approximately 30 cm of caliche and gravel road base was laid for the new road, except in areas where utility work and manholes had not been completed. Grading also occurred in the traffic islands north of the intersection of Ennis Joslin and Alameda Street. Sediments were removed to a depth slightly below the adjacent new road surface, and then backfilled with a gravel and caliche mix. The work was done primarily with a maintainer, with some use of a backhoe in the narrow islands to the north of the intersection. No undisturbed sediments were observed during the island grading.

Bases for Street Light and Utility Poles. Though most of these bases were dug during Phase V (see below), a few were dug while CAR was present during Phase III. The holes for the pole bases were either 0.60 m (2") or

0.90 m (3') in diameter and were dug with an auger to a depth of 1.8 m (6'). Monitors watched this process closely.

Electrical Trenches. Electrical trenches, for laying conduit to connect streetlights and lights in planted areas of the intersection of Ocean Drive and Ennis Joslin, were dug the length of the project area on both sides. With very few exceptions, these trenches were dug into the stabilized sand and gravel that was used to backfill the utility trenches. Monitors observed this digging when the trenches crossed areas that had not been previously undisturbed. All trenches were roughly 0.3 m (1') deep.

A total of 164 features, 11 occupation zones, and six burials were encountered during the first three phases of this monitoring project. Chapter 6 summarizes the features and occupation zones and the information they provide regarding patterns of site occupation.

Phase IV: February 2006

For three days in early February 2006, a CAR archeologist returned to 41NU2 to monitor the digging of holes for the planting of large palm trees, in the new traffic islands. These holes were dug with augurs to a depth of roughly 1.2 m. The backdirt from the augurs, and as much of the profile as could be seen, was observed. No features or burials were encountered.

Phase V (SWCA): March 8-April 24, 2006

After field work by CAR had been completed, TxDOT contracted with SWCA Environmental Consultants, Inc., to monitor excavations of 48 street light pole bases and three shallow trenches for more electrical conduits. With only occasional exceptions, the digging during this project was in sediments already disturbed by the construction monitored by CAR. TxDOT requested that features be designated only if they were considered significant. Three small concentrations of shell and animal bones were observed but were not considered significant. Nothing was collected from the site.

Excavating Human Burials

Whenever human remains were encountered, CAR notified the Project Superintendent and the Director of CAR immediately. The CAR Director in turn notified the TxDOT Environmental Affairs Division in Austin, as required by the established protocol. After an assessment of the nature and context of the remains, CAR was given instructions on how to proceed in each case. Specific excavation details and

circumstances are described in the discussion of the individual interments in Chapter 5. Once excavated, all human remains were returned immediately to the CAR laboratory at The University of Texas at San Antonio where the remains were cleaned and cataloged. During their processing, all remains were kept in a processing facility with limited access. The analysis of the human remains excavated by Ricklis (1997) and reported by Mahoney and Shafer (2003) was integrated into the summary of the human remains recovered by CAR.

Laboratory Methods

The collection from 41NU2 consists largely of samples of feature matrix. With some exceptions, artifacts were not collected. All the cultural materials that were collected were returned to the Center for Archaeological Research. Initial processing of recovered artifacts prior to special analyses consisted of washing, sorting, counting and/or weighing, and storing in archival-quality bags with provenience noted on acid-free paper tags in each bag. Specific analytical methods for special

analyses are described in the Appendixes of special reports and under chapters addressing faunal and skeletal analysis.

All artifacts and faunal remains recovered during Phase I of the monitoring project were returned to the CAR laboratory for processing and analysis (Meissner 2000). All soil samples from features were floated, and the light and heavy fractions were dried. The light fractions were sent for macrobotanical analysis. Artifacts and faunal remains extracted from the heavy fraction were processed and analyzed. The results of the artifact analyses are presented in Chapter 5.

Curation

All cultural materials along with documents generated during field work including the field notes, forms, and photographs were processed for long term curation as directed by guidelines set forth by the Texas Archeological Research Laboratory where this collection is permanently curated. Human skeletal remains also were transferred to TARL.

Chapter 5: Results of Monitoring

Barbara M. Meissner

Twenty-one major trenches, and several smaller trenches, were dug during this project. We collected data on 164 features exposed in these trenches, identified 11 occupation zones, recovered seven human burials, and collected 125 artifacts. Fifty-eight radiocarbon dates were obtained during the project, including 12 humate dates associated with coring activities, and 46 dates on charcoal, organic material, and shell associated with features or occupation zones. This chapter describes the trenches, the human burials, and the small collection of artifacts. The subsequent chapter provides information on feature associated radiocarbon dates, features, and the 11 occupation zones identified.

Monitoring Overview

It is a truism that archeologists often feel as if they spend all their time in the field writing notes and taking pictures—only to complain later, when they return to the lab to write up the project, that they did not take enough notes and pictures. This was certainly true in this case. When this project began, no one expected it to last as long as it did or for the site to be as complex as it was. This chapter is based on notes and photos taken by a number of people over a period of seven years, during which the period of intense daily activity lasted two and a half years. There were two primary Project Archeologists (PAs) during the project, Cynthia L. Tennis during Phase I, and Barbara A. Meissner for the remainder of the project. In addition, over the course of the project there were short periods of time when other qualified staff members of the CAR were required to fill in. These instances were infrequent and of brief duration.

Unfortunately, some data have been lost completely. On the night of May 12, 2004, the CAR vehicle, parked at the motel where CAR personnel were staying, was burglarized and a number of items, including tools and a camera, were stolen. A plastic case containing profiles, feature maps, and, most importantly, the field notebook containing notes on the project dating from March 26, 2003, to November 10, 2003, was taken. CAR notified the TxDOT Environmental Affairs Division of the event. Although this loss of data was disastrous, it could have been a great deal worse, as much of the data in the notebook was also listed on Feature forms, the Field Sack Log, and on the trench maps, all of which were not in the vehicle. Project photographs and the weekly reports also provide a source of backup information. Although detailed descriptions of daily activities between late March and early November, 2003, were lost, information on what

cultural deposits were observed was recoverable from these other sources.

Between August 15, 2000, and November 3, 2005, CAR monitored construction activities associated with road and utility improvements in the right-of-way of Spur 3 within the site boundary of 41NU2. Major improvements included: 1) rebuilding and widening Ennis-Joslin; 2) ending Alameda Street at a T-intersection with Ennis-Joslin; 3) rebuilding the portion of Alameda Street that would now become Ennis Joslin; 4) simplifying the Ocean Drive/Ennis-Joslin intersection; and 5) replacing storm and sanitary sewers, water, and gas lines in the area and remove old utility lines. Monitoring of large tree plantings took place in February 2006. Monitoring of the placement of some additional street light footers and shallow trenches for electrical conduits that connected them was carried out by SWCA Environmental Consultants, Inc. The results of that project are included at the end of this chapter. And finally, once in June and again in August 2007, CAR staff was also present to help extract the soil cores used for a number of the special studies in this report. Detailed project maps, showing the location of all monitored excavations, including all new utility trenches, utility locations pits, and the pits dug to either place or remove manhole boxes are presented in Figures 5-1 through 5-5. The figures (5-1, 5-2 and 5-4) also show the locations of the three soil cores extracted by the project geomorphologist.

The understanding of the nature of the site, its geomorphology and site formation processes, evolved over time. In the early stages of the project, the complexity of the subsurface sediments was not entirely recognized. Similarly, the spatial relationship between the two dunes and the ridge of Pleistocene deltaic Beaumont Clay Formation (the Beaumont) was not entirely comprehended until sufficient numbers of trenches were excavated along the project ROW to allow us to piece together the spatially isolated pictures into a composite contiguous image. As a result, some of the information presented in this report differs from interpretations provided in the four interim reports (Meissner 2000, 2003, 2004, 2005) previously submitted to TxDOT-ENV. The differences are a natural product of the evolution of our understanding of the site and the landform upon which it sits.

Figure 5-6 shows a schematic profile of the site, looking west from the False Oso, showing the complex nature of the geomorphology. Note that the horizontal scale in Figure 5-6 is 20 times the vertical scale. This is necessary to present the

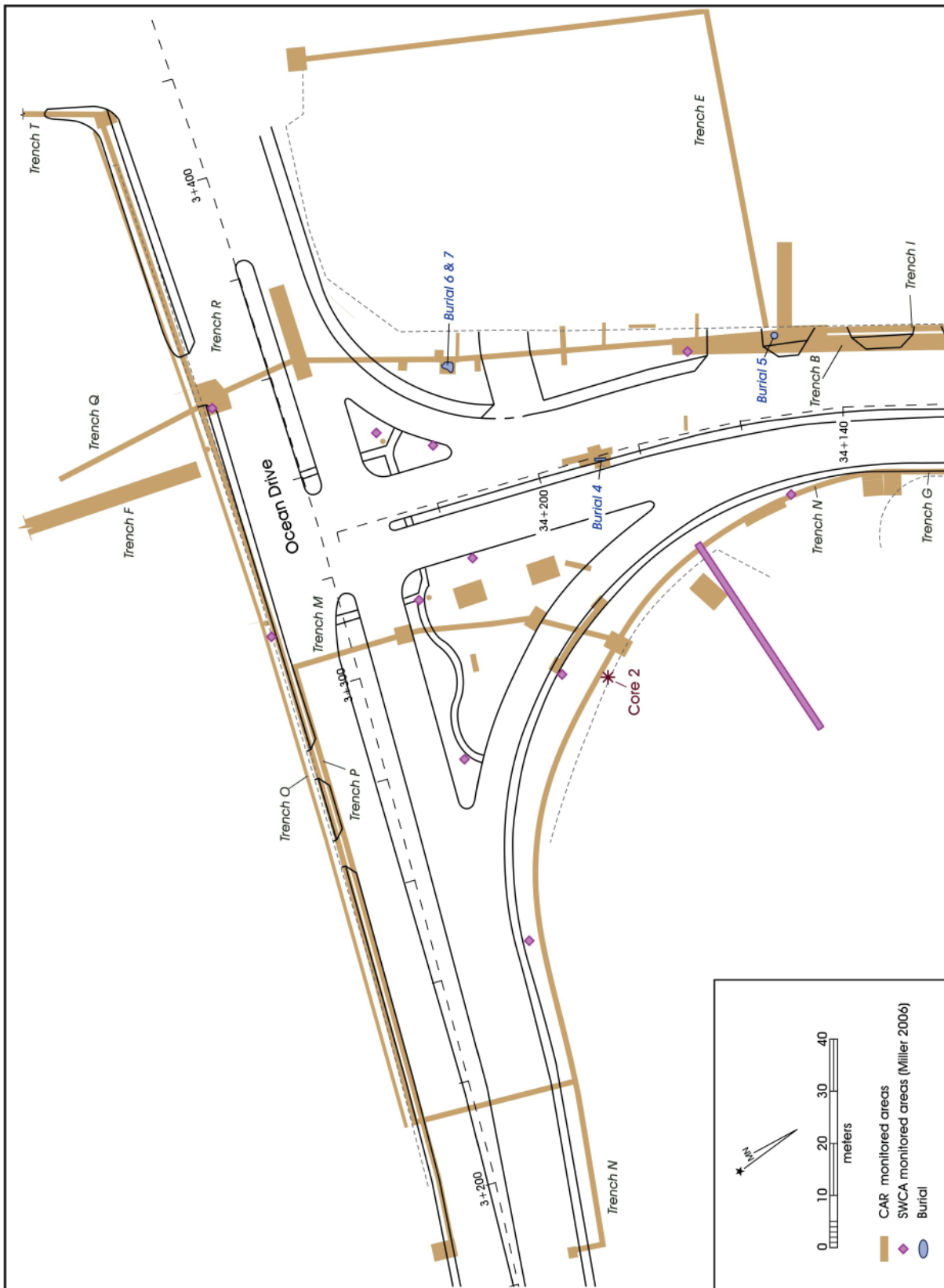


Figure 5-1. Detail of project map between 34+120 and 34+325 (including all work on Ocean Dr.) showing locations of all monitored excavations, burials, and sediment cores.

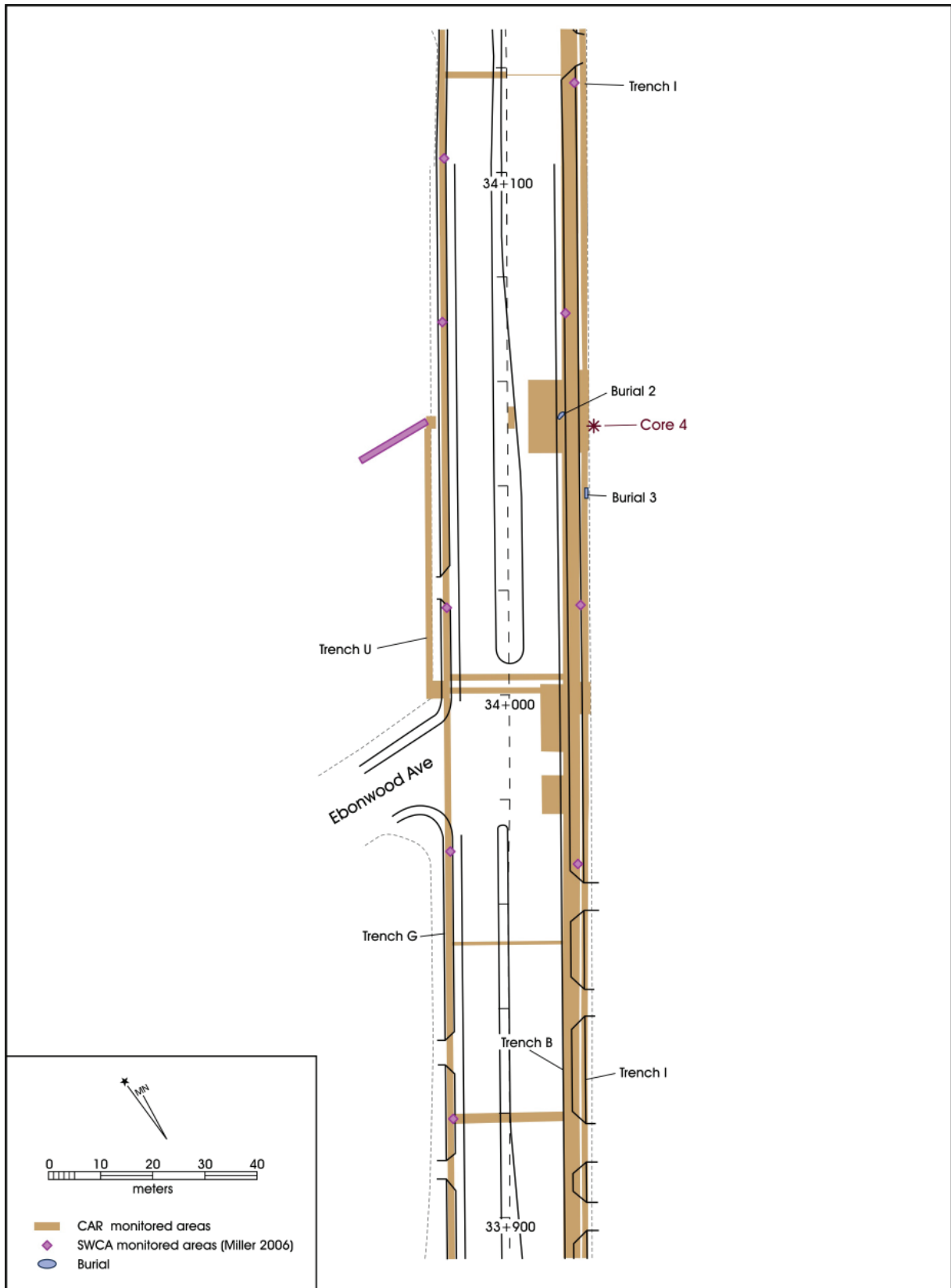


Figure 5-2. Detail of project map between 33+900 and 34+120, showing locations of all monitored excavations, burials and sediment cores.

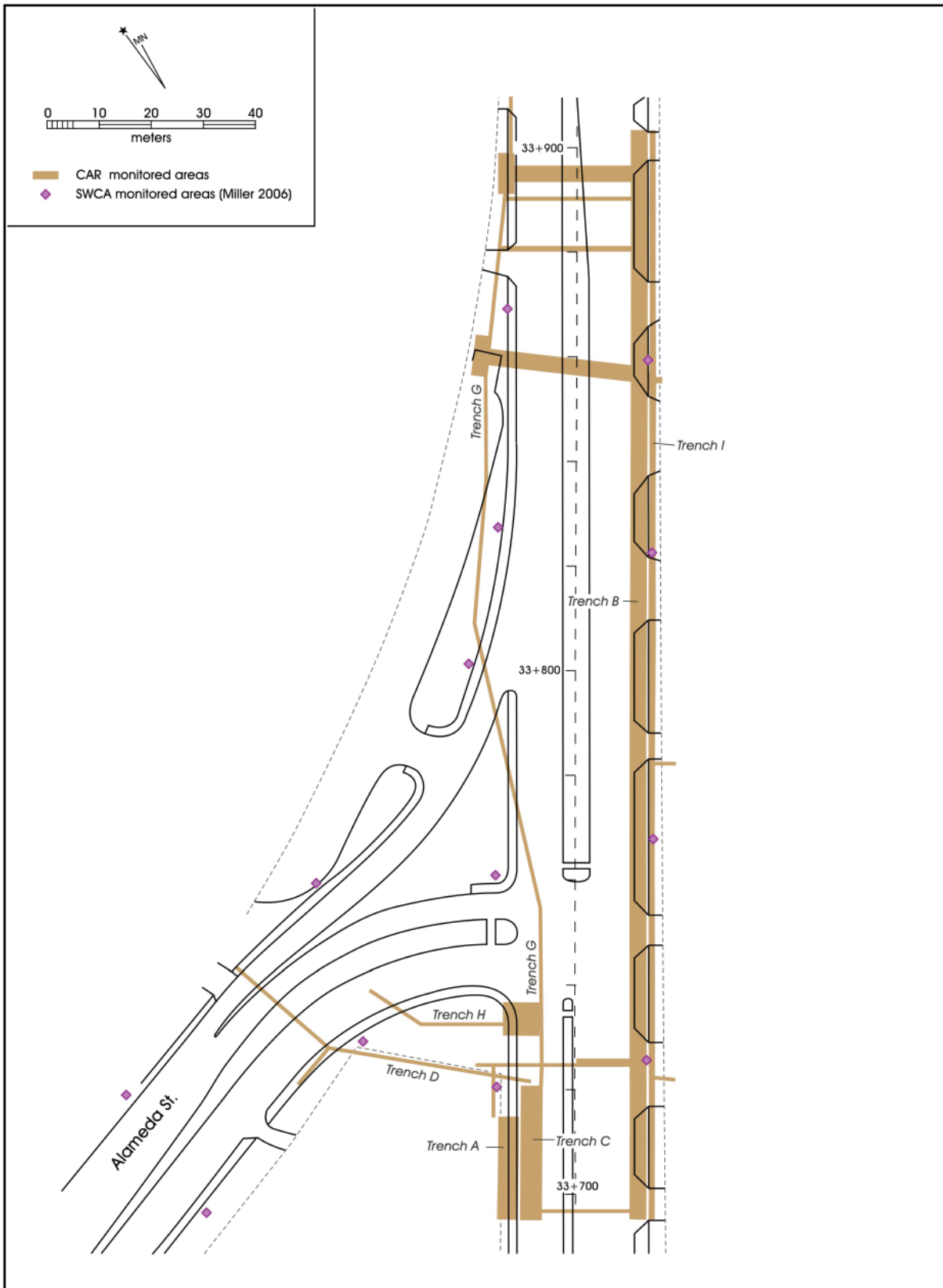


Figure 5-3. Detail of project map between 34+120 and 33+700, showing locations of all monitored excavations.

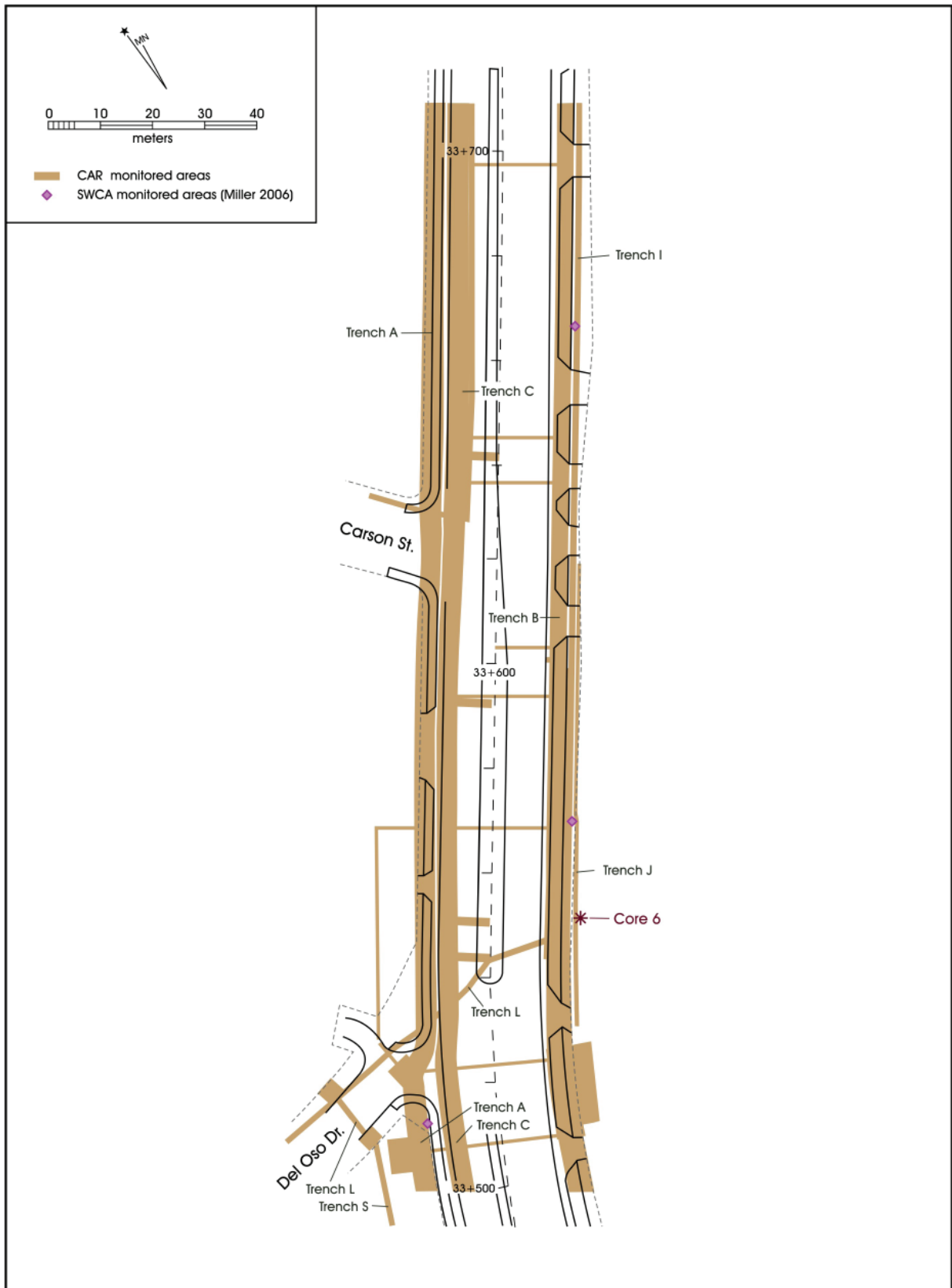


Figure 5-4. Detail of project map between 33+500 and 33+700, showing locations of all monitored excavations and sediment cores.

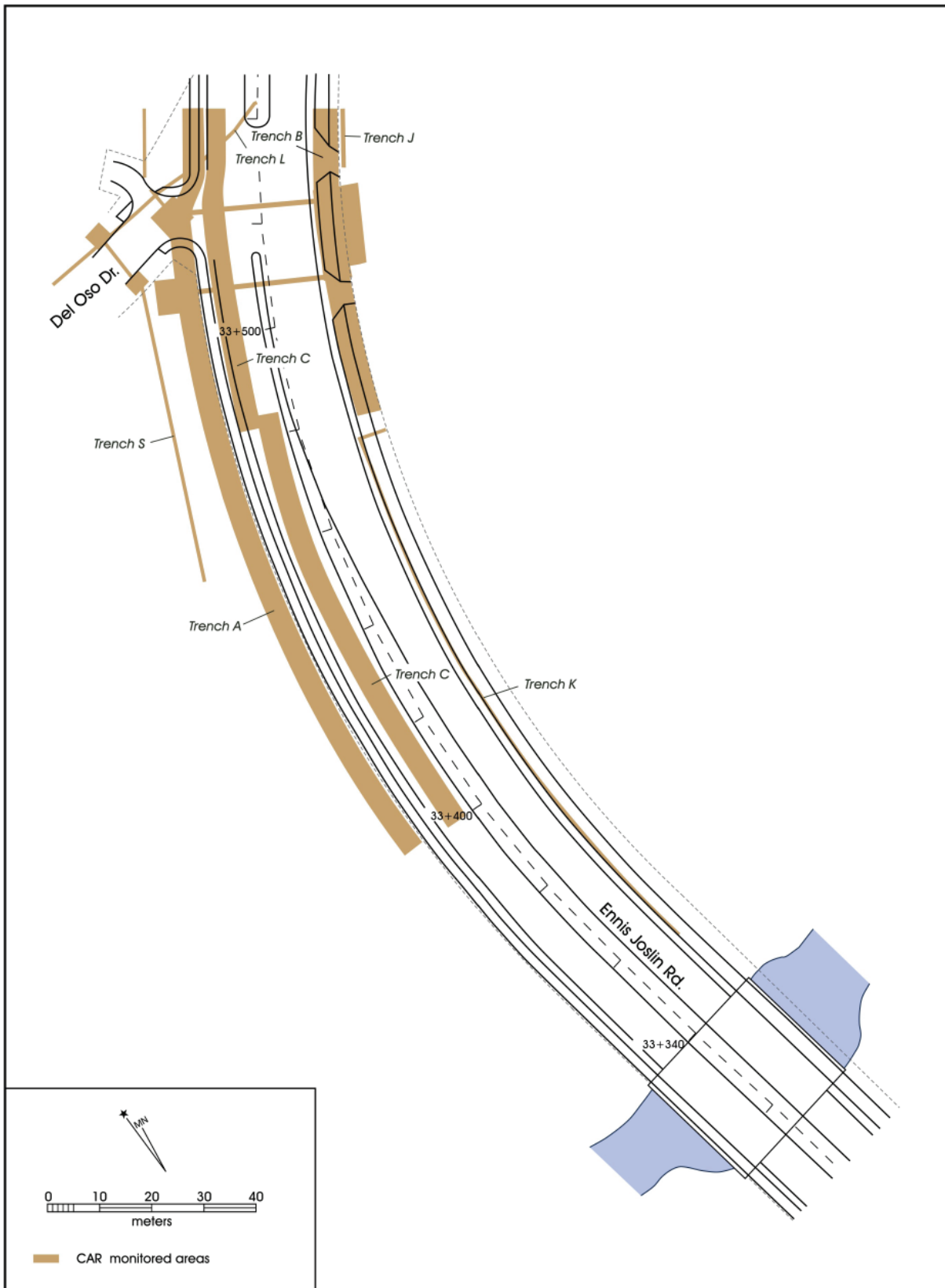


Figure 5-5. Detail of project map between 33+320 and 33+500, showing locations of all monitored excavations.

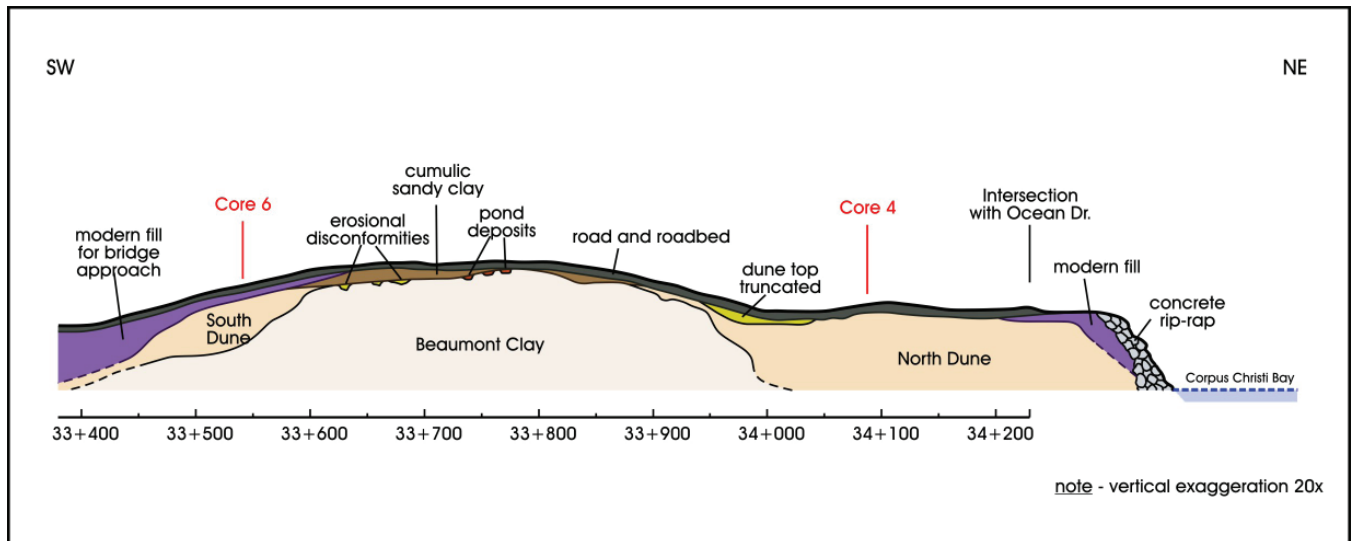


Figure 5-6. Schematic profile of NU2, looking west from the False Oso mudflats, showing the Beaumont Ridge with dune deposits to the north and south. Note that horizontal scale is 20 times the vertical scale, thus grossly overemphasizing elevation changes. Based on profiles of Trenches B, I, and F.

north/south extent of the site on a single page, but results in a gross overemphasis of elevation changes. Utility trenches and other modern disturbances are not shown, and the location of erosional channels and depressions filled with pond deposits on the Beaumont ridge are approximate.

Figure 5-7 is a schematic profile running east/west along Ocean Drive roughly 80 m south of the bluff above Corpus Christi Bay. Note that the ridge of Beaumont to the west and the adjacent North Dune forms part of the depression in which the large ephemeral pond seen in old maps and photographs of the site developed (e. g. Figure 5-8). No information is available to estimate how far the North and South Dunes extend to the west except at the northern-most part of the site.

Although other excavations were monitored during the project, trenches dug to place new utilities were considered the most likely places in which human burials or other important cultural remains could be encountered. Therefore, we begin the discussion of the findings of the monitoring with the descriptions of the trenches.

Trenches

Trench A

Trench A, was excavated entirely during Phase I of the project between August-October 2000. This trench, which was 4 m wide, was located on the west side of Ennis Joslin, beginning

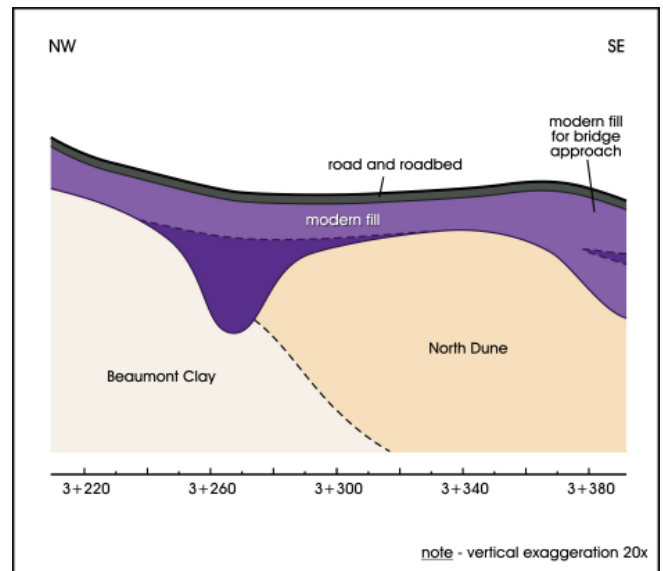


Figure 5-7. Schematic profile of the northern part of NU2, along Ocean Dr showing Beaumont clay ridge to west, the pond area, and the slope down to the False Oso, near its mouth at Corpus Christi Bay. Note that horizontal scale is 20 times the vertical scale, thus grossly overemphasizing elevation changes. Based on profiles in Trenches M and R.

just south of the intersection with Alameda Street. The center of the trench was 12 m west of the project centerline (CL) at station 33+714 (Figures 5-3 to 5-5). It continued south, with two interruptions, moving away from the CL, to 33+406 (Figure 5-5). At that point, the center of the trench was roughly 18 m from the CL. The trench varied in depth from 80 cmbs to 260 cmbs.

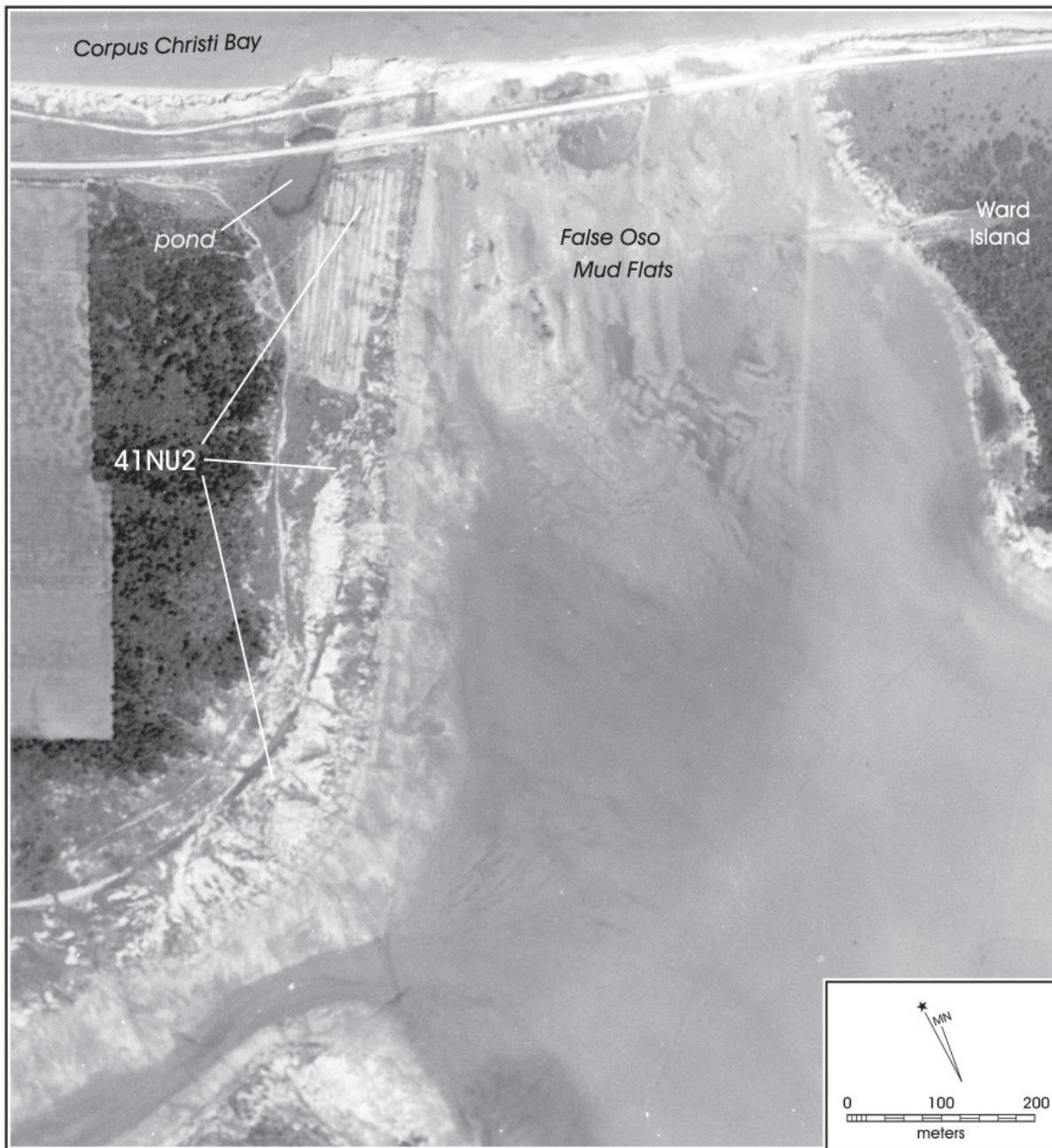


Figure 5-8. Aerial photograph of 41NU2 taken in 1936, showing 41NU2 before urban development. Note the pond near Corpus Christi Bay.

Trench A provided the first view of the way in which the simple stratigraphy at the center of the site noted by Ricklis (1997:20-21, 31) changed to the “striped” appearance typical of clay dune deposits (Huffman and Price 1949:118; Ricklis 1997:37; see also Figure 2-1). Profile 1 in Figure 5-9 shows a typical profile of the South Dune. Four paleosols are evident, separated by thick layers of light colored clay. Unfortunately, the exact location in Trench A at which the dune deposits end, replaced by the rising Beaumont Ridge, was destroyed by previous utility trenches.

As Trench A continued north (Profile 2 in Figure 5-9), the sediments become relatively shallow clay dune deposits sitting above the Beaumont clay, which has dropped about 50 cm in the 200 m between the two locations profiled. The banding of dark and light sediments is beginning to fade, and by the time Trench A reaches Profile 3 in Figure 5-9, near the center of the site, the strata have become much simplified, showing only a dark brown sandy or silty clay loam that has accumulated on top of the remnant of the Beaumont Clay Formation we are calling the Beaumont Ridge (see Chapter 2), with no sign

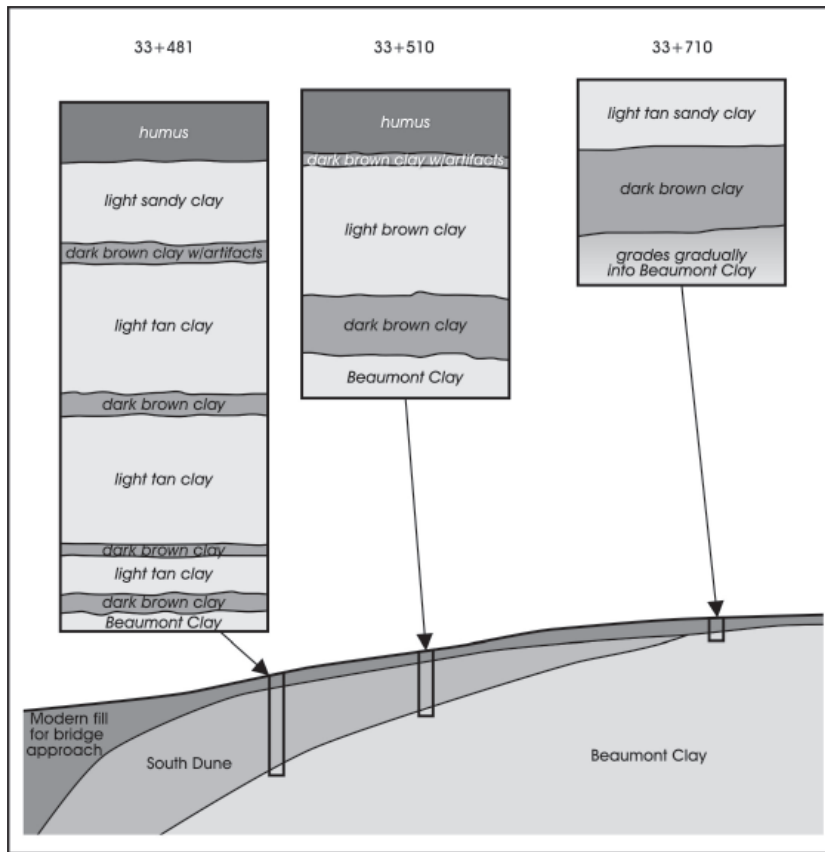


Figure 5-9. A schematic profile of the South Dune and Beaumont Ridge, looking west, showing profiles from Trench A at three locations.

of dune deposits. In some areas near the highest point on the current landscape within the site boundaries, the Beaumont clay actually lies directly beneath the old road bed, presumably because the cumulic deposits were removed during the initial construction of Ennis Joslin and S. Alameda Street.

When completed, Trench A crossed two of the four subsurface zones of the site, the South Dune and the Beaumont Ridge (Figures 5-6 and 5-7). Nine features were recorded in this trench (Features 1-5, 7, and 9-11). Their locations can be seen in Figures 5-3 to 5-5.

Trench B

The first part of Trench B was excavated during Phase I (Figure 5-4). This section ran from 33+525 north to 33+595. The trench was 2.4 m wide with the center about 11.5 m east of the CL. During this early part of the excavation of Trench B, five features were encountered (Figure 5-4). Trench B was discontinued when TxDOT decided that there were too many existing utilities for which they had no maps. CAR's part of the Spur 3 project was put on hold while new plans for the monitoring methodology and the actual construction plans themselves were considered.

During Phase III, beginning in April 2003, the remainder of the trench was excavated. In the interim, some construction plans had been changed, resulting in a change in the location of Trench B to center at 12.0 m east of the CL, and widening the trench to 3 m (Figure 5-4). This required the re-excavation of part of the trench that had been dug in 2000. The depth of the sterile Beaumont varied a great deal on the east side of the road, as it did on the west, from immediately below the roadbed between about 33+720 to 33+800 to an unknown depth more than 240 cm below the ground surface in the areas south of 33+440, and more than 3 m below the ground surface north of 34+003 (Figure 5-6).

Excavation of Trench B provided a view of the subsurface of almost the entire site in a line, running roughly north and south, about 73 to 105 m west of the currently stabilized bluff above the False Oso mud flats. This resulted in a much better understanding of the structure of the site from south to north (Figure 5-6).

It was evident from Trench B that the remnant ridge of Beaumont clay, seen by Ricklis (1997) and also present in Trench A, had separate dune deposits built up on both the southern and northern faces as seen in Figure 5-6. Between about 33+600 and 33+880 in Trench B, the top of the Beaumont clay is near the ground surface. Both south and north of this area, the upper surface of the Beaumont clay slopes down and is covered with Holocene clay dune deposits (Figure 5-6).

Through the South Dune

Figure 5-10 is a profile of the east wall of Trench B at 33+566-567, about 5 m north of Ricklis' Test Unit 5 (Ricklis 1997:25). Examination of Ricklis' profile of Unit 5 (Ricklis 1997: 28) shows almost identical layers of dark and light sediment, although the uppermost paleosol is darker in Figure 5-10. Note that the lowest level in both profiles (Ricklis 1997:28 and Figure 5-10) appears to be a continuation of OZ-G, found in a 30 cm thick paleosol in Trench I (see Trench I section below). The top of this 30 cm paleosol was observed to drop from 80 cm bgs at 33+600 to 110 cm bgs at 33+584, where it could readily be seen to continue south, below the level of the trench bottom. It is not unreasonable, therefore to tentatively equate the lowest paleosol in Figure 5-10, whose top is at about 120 cm bgs, whose thickness is an unusual 30 cm,

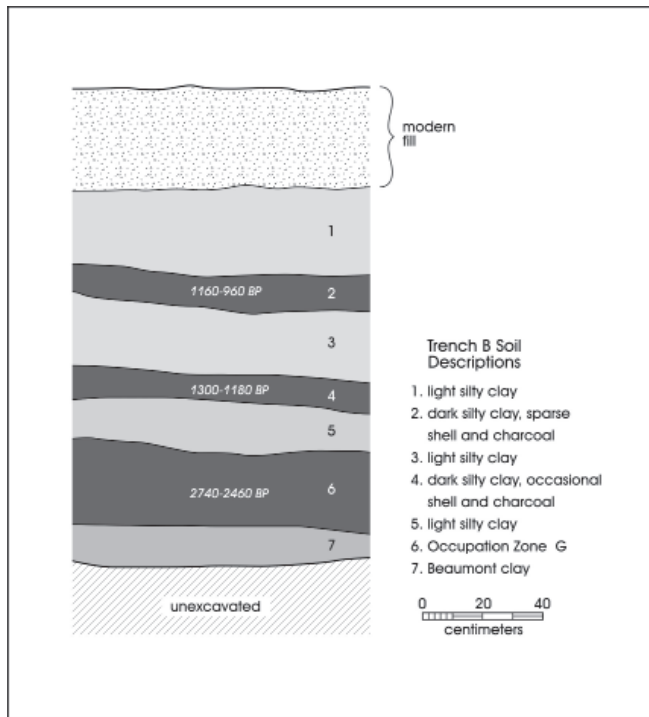


Figure 5-10. Profile of the east side of the South Dune in Trench B at 33+566-567, showing estimated dates for three paleosols.

and which is full of cultural debris such as shell, charcoal, and burned clay, with OZ-G, as seen in Trench I. A detailed discussion of OZ-G is provided in Chapter 6.

Through the Beaumont Ridge

By about 33+605, the thin remnants of OZ-G are lying directly on top of the Beaumont clay surface, beneath the dark brown cumulic sandy clay that overlies the Beaumont were there is no dune deposit. Because of unfortunately placed old utilities, this is the only place on the site where the upper edges of dune deposits can be observed in contact with the Beaumont clay.

On the east side of the site, between about 33+607 and 33+980, the Beaumont lies near the ground surface, with varying amounts of cumulic fine sandy clays above it. Naturally refilled erosional channels were seen most often in the Beaumont Ridge area (Figure 5-11).

For the most part, the dark sediments above the Beaumont, when present, were sterile of cultural remains except for an occasional shell. However, an occasional scatter of shell and charcoal or distinct feature was located, always lying either directly on or just above the Beaumont clay surface. The largest such feature, Feature 57, was located in both Trenches B and I at about 33+716. This feature lay directly on top of the Beaumont clay. Another interesting feature, which will be discussed in Chapter 6 is Feature 54, a large pit hearth, was dug into the Beaumont clay.

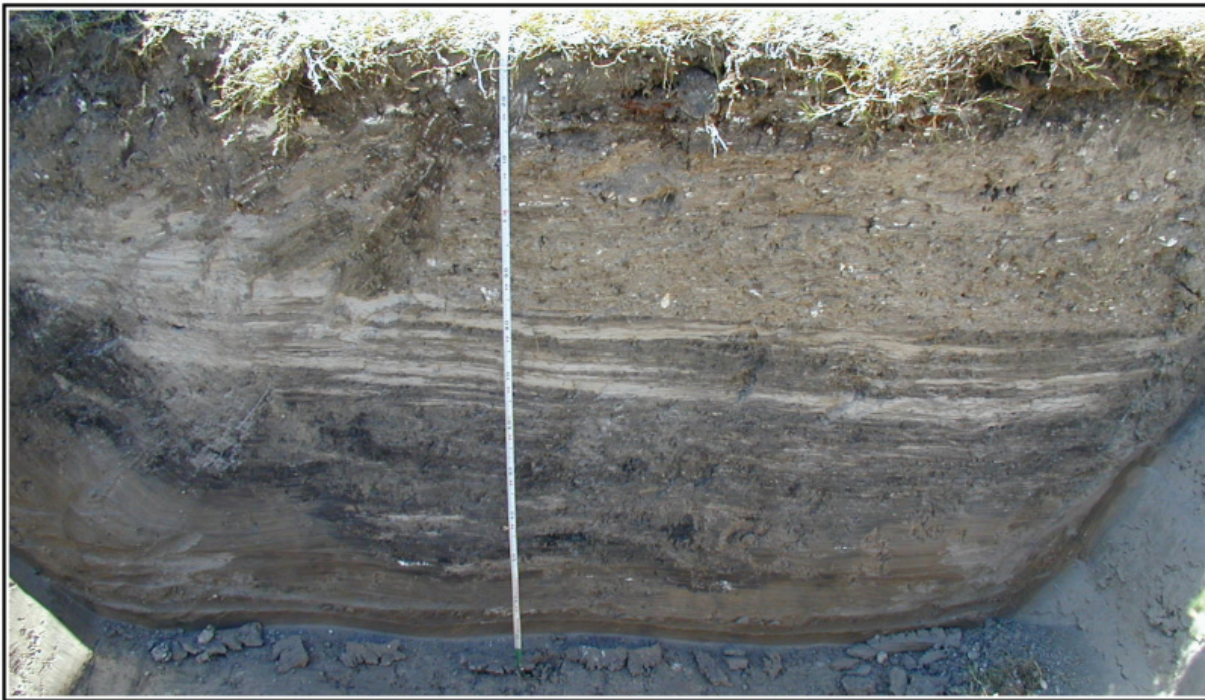


Figure 5-11. Photographs of Beaumont Ridge zone at 33+689, showing large, naturally refilled erosional channel. Note shell at bottom of erosional channel and modern fill above the natural laminated sand and clay fill.

Through the North Dune

The top of the Beaumont clay began to drop in elevation around 33+800 and by 33+920 the edge of another series of dune deposits became noticeable. In a profile drawn of the west wall at 33+923-925 (Profile 1 in Figure 5-12), the early indications of the North Dune are becoming apparent, and the north wall of a lateral immediately west of Trench B at 33+985 (Profile 2 in Figure 5-12) shows the increasing complexity of the dune deposits and the increasing depth of the Beaumont clay.

Only 11 m further north, in a profile of the west wall of a large pit dug to place a large manhole box (Profile 3 in Figure 5-12), the complexity of the southern part of the North Dune becomes more apparent. It should not be assumed, however that the complex strata shown in Profile 3 continue more or less in this fashion to the bluff above Corpus Christi Bay. Unlike the South Dune which displayed a relatively simple series of three to four paleosols 10 cm or thicker, the stratigraphy of the North Dune varied over its length and width in Trench B, and the variation continued passed Trench B (see Trench F, below).

This variation was manifested largely in the darkness and thickness of paleosols. None of these were as dark or as thick as those recorded in the south dune. They could vary a great deal over a relatively short distance, becoming darker and thicker, then becoming so thin and so light that they would be almost impossible to separate from light clays, then becoming darker and thicker again (Figure 5-13).

As Trench B continued north across the north dune, layers of sediment began to slope upwards, as can be seen in Figure 5-6, a trend that continued until about 34+060, after which the layers of sediment tended to slope slightly downwards to the north, towards Corpus Christi Bay (Figure 5-6). Throughout the North Dune seen in Trench B, the sediments also tended to slope down to the east, towards the False Oso mudflats.

As Trench B continued north, the contrast between the dark and light layers below about

1 m under the road surface decreased somewhat, and many of the thinner dark layers faded out almost completely. As Figure 5-13 shows, the amount of change that can take place over even the approximately 3 m between the center wall of Trench B (about 12 m east of the centerline, shown in

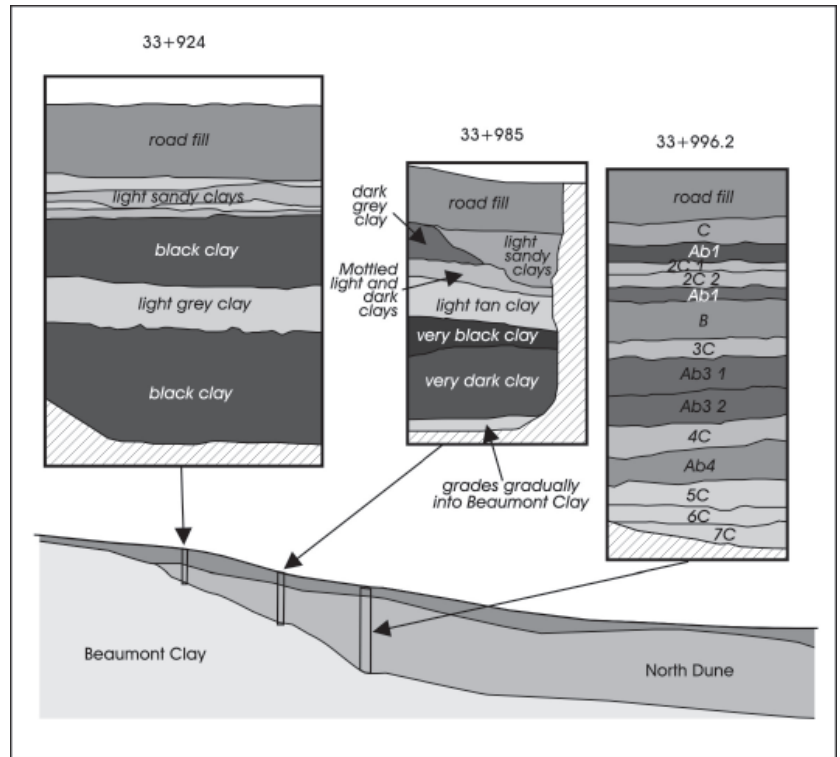


Figure 5-12. A schematic profile of the North Dune and Beaumont Ridge, looking west, showing profiles from three locations in Trench B.



Figure 5-13. Photograph showing Trench B at approximately 34+080, as the first half of a new section has been completed, showing the wall at the center of the trench (foreground) and the wall at the western edge of the trench (background). Looking southwest.

the foreground of the photograph), to the west wall (about 10.5 m east of the centerline, shown in the upper part of the photograph) is remarkable.

In this area of the North Dune, the two upper-most paleosols were only about 10-15 cm apart at the center of the trench but most of the upper paleosol has been removed by road construction in the western half of the trench and, since OZ-E does not extend to the west wall, the second paleosol is only faintly visible (Figure 5-13). These two paleosols continue to get closer to each other as they progress east, so much so that in Trench I, adjacent to the east of Trench B, it was difficult to separate them during excavation, given the technique in use.

Beginning at roughly 34+000, with evidence seen in the Ab1 layer in Greaves' (2004) figure reproduced in Profile 3 of Figure 5-12, there was an increasing scatter of charcoal fragments, tiny burned clay fragments and shell fragments in the uppermost paleosol. This was merely noted at first, but when this scatter became denser and persisted for 42 m, it was designated an occupation zone. Examination of notes taken previously indicated that a similar phenomenon had been noted in Trench C (see below). This first OZ was designated OZ-A, while the first OZ in Trench B was designated OZ-B. Three more OZs were identified in Trench B, OZs -D, -E, and -F. They will be discussed in the following chapter.

In December 2003, a single human burial, Burial 2, was located in Trench B, with the top of the skull found 10 cm below the bottom of the road bed, 10.63 m east of the centerline at 34+053. This burial will be discussed in greater detail in a later section of this chapter.

Trench B ended at 34+170, however a large percentage of the last 15 meters was disturbed by previous utility trenches. Once the clearing had proceeded nearly to the north end of the trench, Trench B was re-excavated and several utility lines, including a storm sewer, water, gas, and sanitary sewer lines were installed. CAR personnel monitored this re-excavation, as previously undisturbed sediments were sometimes impacted in the process of re-digging this trench, especially when widening it for manhole boxes, etc. Later, the curb of the new road and a sidewalk were installed above the trench.

Sixty-six features was identified within Trench B (Figures 5-1 through 5-5), including the five features discovered during the initial excavations of the trench in 2000 (Meissner 2000). These features are described in the following chapter.

Trenches C, G, and N

Trenches C, G, and N are in reality a single trench dug at different times. The trench sections were given different designations in the field and in the interim reports, but for this final report, they will be described together. They form a western version of Trench B, and provide a view of the west side of Ennis Joslin from 33+400 north up Ennis Joslin and old S. Alameda Street with a turn to the west just south of Ocean Drive (Figures 5-1 through 5-5).

Trench C

Trench C was excavated in the earliest part of Phase III of the project, after changes in utility construction plans moved the location of a large storm sewer, originally intended to go into Trench A, 5 meters east. Trench C was centered about 7 meters from the centerline, east of Trench A and extended from 33+725 south to 33+400, with a gap at the Del Oso intersection (see Figure 5-4). Trench C was 2.2 m wide and varied in depth from 1 m to 2.9 m, depending on the proposed grade of the storm sewer and the depth of the Beaumont clay. Beginning at about 33+490, a layer of modern fill was observed above typical clay dune sediments. As the trench continued south, the surface of undisturbed dune sediments dropped lower in the profile and the percentage of fill increased until, at 33+460, the entire profile was modern fill. The trench cut through modern fill until it ended at 33+400.

The first of eleven OZs was located in this trench. At the time it was simply described as an area within which a thin layer of highly scattered shell, charcoal, and animal bones extended for roughly 18 m in the trench between 33+558 and 33+576 and at a depth of 70-80 cm. This OZ contained a number of areas in which the shell scatter was denser, however only one such concentration was designated as a feature at the time, Feature 17. It was not until the OZs in Trench B were discovered that previous notes were reexamined and OZ-A was designated, based on the description (Meissner 2004).

Trench G

Trench G was a northern extension of Trench C (see Figure 5-3), on the west side of Spur 3. Unlike Trench C, Trench G was only 1 m wide. It was approximately 220 cm deep. At first, it extended across the empty lot at the junction of Alameda Street and Ennis Joslin Road and then turned north, crossed several traffic islands and ended in the middle of the right lane of Alameda (Figure 5-3). This was an area at the top of the Beaumont clay ridge, and no dune deposits or cultural features were noted. In one area between roughly 33+755 and 33+758, however, a layer of extremely black, glossy, highly

organic clay, truncated at the top by previous construction, was noted. It is believed that this deposit represents the remnants of a small ephemeral pond (see Figure 5-6), similar to the one shown on old maps of the site near Ocean Drive (see Figure 5-8). This deposit was only about 30 cm deep in Trench G (see Trench H).

Trench G was later extended, continuing north under the new sidewalk on the west side of Ennis Joslin to a large manhole box at 34+132, beyond which it was designated as Trench N (Figures 5-1 to 5-3). During most of its length, this section was dug to ca. 215 cm bgs. However, from 34+102 to the end, it was only dug 155 cm deep. The southern part of this extension of Trench G was located on the Beaumont Clay ridge and no features or other signs of cultural materials were noted when sediments were undisturbed. North Dune sediments

did not appear in Trench G until approximately 33+990-994 (Figure 5-14). At 34+012, the uppermost dark paleosol began to contain a thin layer of shells and shell fragments, scattered charcoal and small burned clay nodules and scatters of fish and mammal bone. As Trench G continued north, it became clear that this was another OZ. It was designated OZ-J (Figure 5-15). It extended to ca. 34+120 in Trench G, beyond which most of Trench G appeared to be dug through previous disturbances. OZ-J contained ten Features: 148-152, 154-156, 159, and 161, all but the latter of which were in Trench G. Three additional features were encountered in Trench G: one (Feature 153) was above OZ-J and two (Features 157 and 158) were below OZ-J (see Chapter 6) for a total of twelve features identified in this trench. All of the features in OZ-J were truncated to some extent on the eastern side of the trench, due to construction of the original sidewalk.

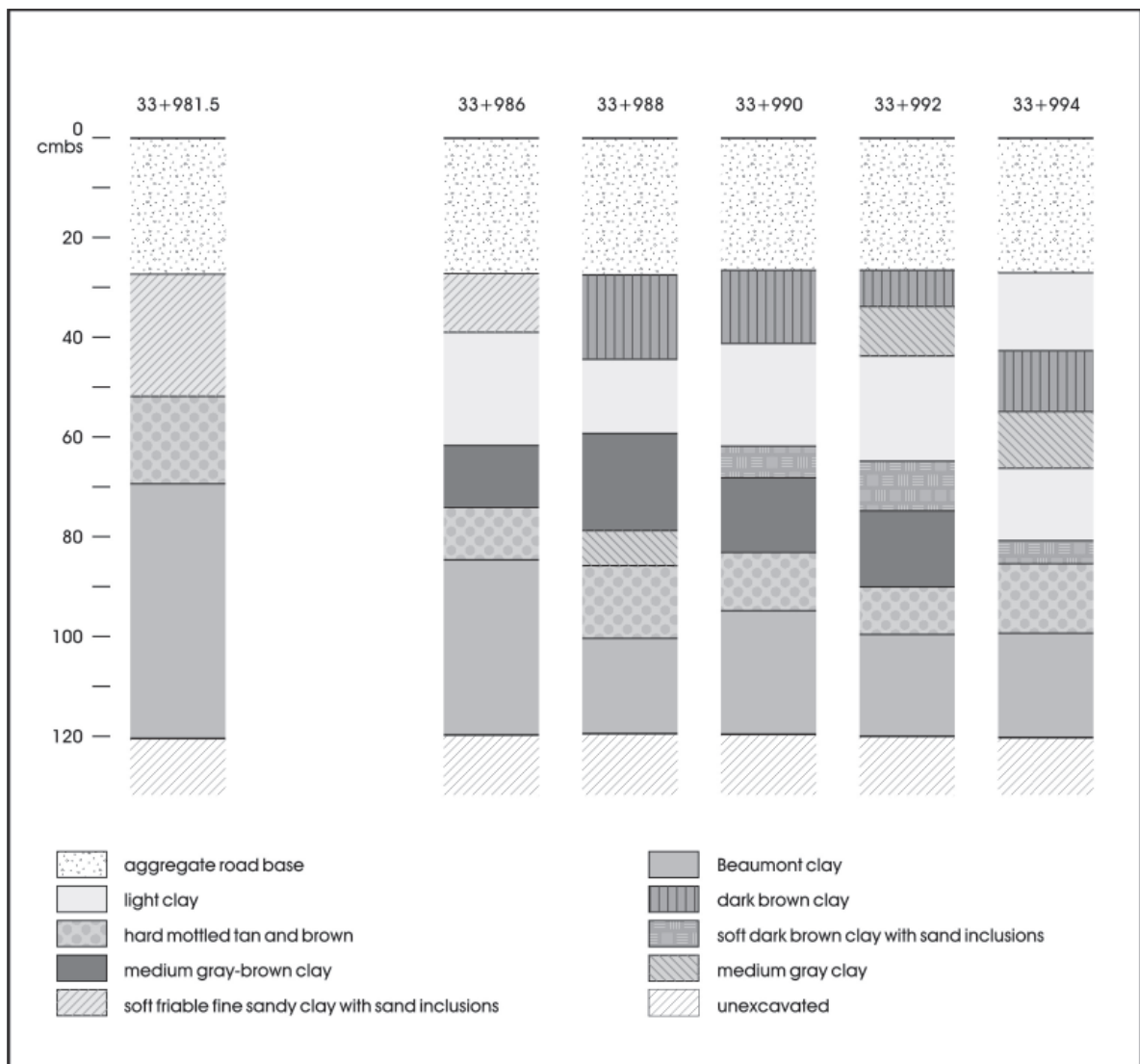


Figure 5-14. A series of 20-cm strip profiles of the west wall of Trench G at the beginning of the North Dune. Note: First gap is 4.5 m due to a disturbance, all other gaps are 2 m.

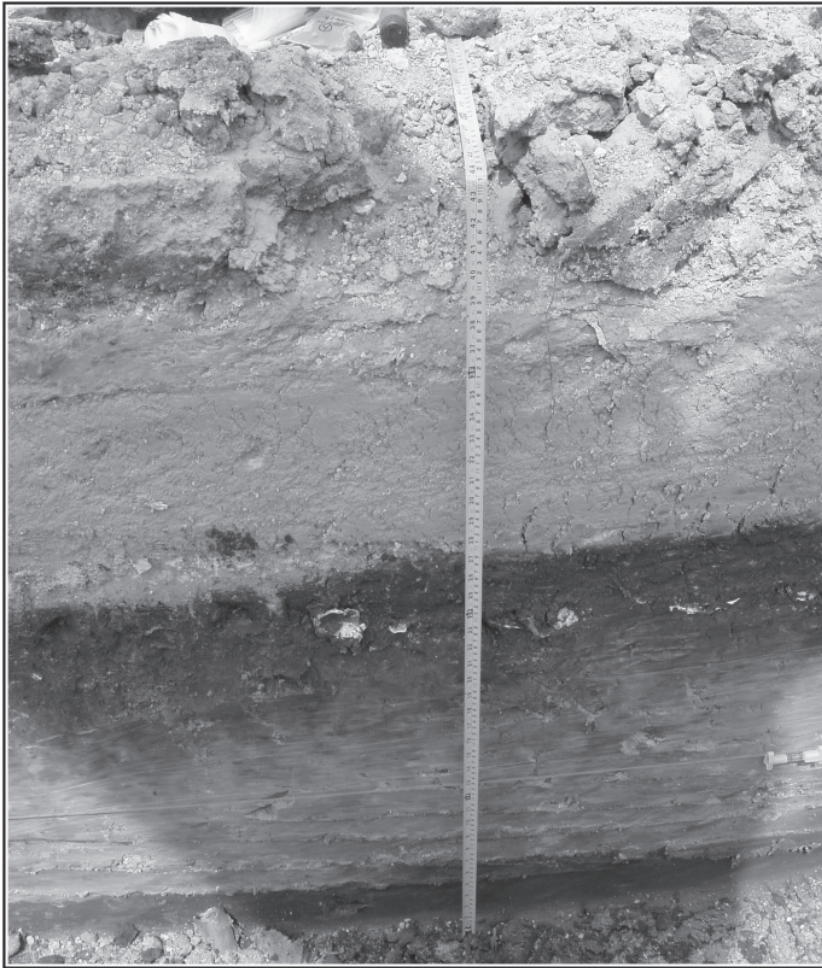


Figure 5-15. A photograph of the west wall of Trench G, showing Feature 149 in OZ J.

Trench N

Trench N is a continuation of Trench G, beginning at the manhole box at 34+132 and continuing around the curve from the west side of Ennis Joslin to the south side of Ocean Drive (Figure 5-1). The southern-most part of Trench N was in the northern dune deposit, but as the trench continued north to about 34+179, the extremely black pond deposits were seen above the dune deposits, which dropped sharply until they were no longer seen in the trench (Figure 5-16). These pond deposits continued as Trench N made the curve to the west, ending at about 3+260 on Ocean Drive (see Figure 5-1). However, the northern dune deposits did not return. Instead, Beaumont Clay emerged from beneath the black pond deposits, with dark loamy sediment developing above the Beaumont (Figure 5-17). The pond deposits became very shallow and disappeared as the trench continued to the west. A thin layer of shells and flecks of charcoal was observed in the dark loam above the Beaumont clay, especially between about 3+260 and 3+257 but was

too scattered to define as a feature (Figure 5-17). It was later determined that this thin layer of shells was probably part of OZ-K (see Chapter 6).

Between 3+232 and 3+235, Feature 147 was encountered. This feature consisted of three small (30-45 cm) pits, one of which was dug into the Beaumont Clay beneath the feature. The pits were densely packed with Oyster (*Crassostrea virginica*) and Lightning Whelk (*Busycon perversum*), and surrounded by a dense scatter of shell, fish and mammal bone, charcoal, and small burned clay nodules. Feature 160 was a small concentration of shell about 10 cm above the Beaumont Clay.

By about 3+210, the Beaumont Clay was beneath just 10 cm of topsoil. Although the surface of the Beaumont Clay dropped slightly as the trench continued to the west along Ocean Drive, the dark loamy clay above it contained no more than an occasional shell west of Feature 160. Features 147 and 160 were the only feature identified in Trench N. The trench ended at 3+185.

Taken together, Trenches C, G, and N had fourteen features and two OZs. This is substantially less than the roughly equivalent Trench B on the east side of the road.

Trench D

Trench D was excavated during Phase II of the project. It began on the west side of Ennis Joslin Road at 33+724 and continued west to the curb of the south edge of S. Alameda Street where it turned northwest to cross the street (see Figure 5-3). The trench was 60 cm wide and was dug to a depth of 120 cm. Below about 20 cm, the sediments in the trench were undisturbed until the edge of Alameda Street was reached, where old utilities, most of which were no longer functional, had been installed. Other than an occasional shell, there were no cultural materials observed. Near the eastern end of the trench, the sediments above the Beaumont included a thin (ca. 10-15 cm) layer of unusually dark highly organic clay. Later excavations to the north of Trench D (see Trench H, below) showed more of this sediment that developed in an area where rainwater collected in depressions in the impermeable Beaumont clay and formed small, ephemeral ponds (see Chapter 2).

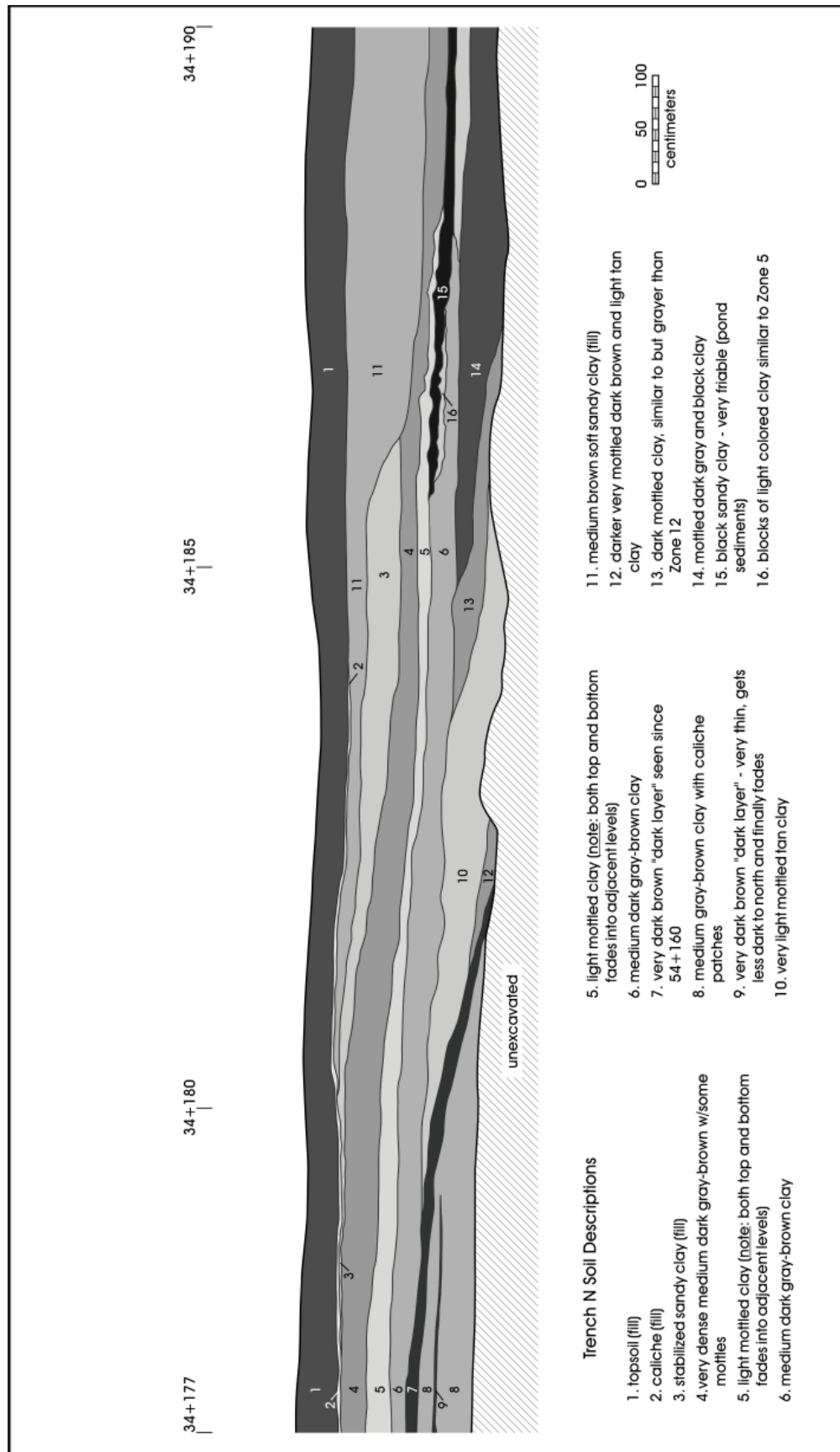


Figure 5-16. Profile of the west wall of Trench N, showing north dune deposits ending and pond deposits beginning, between 34+177 and 34+190.

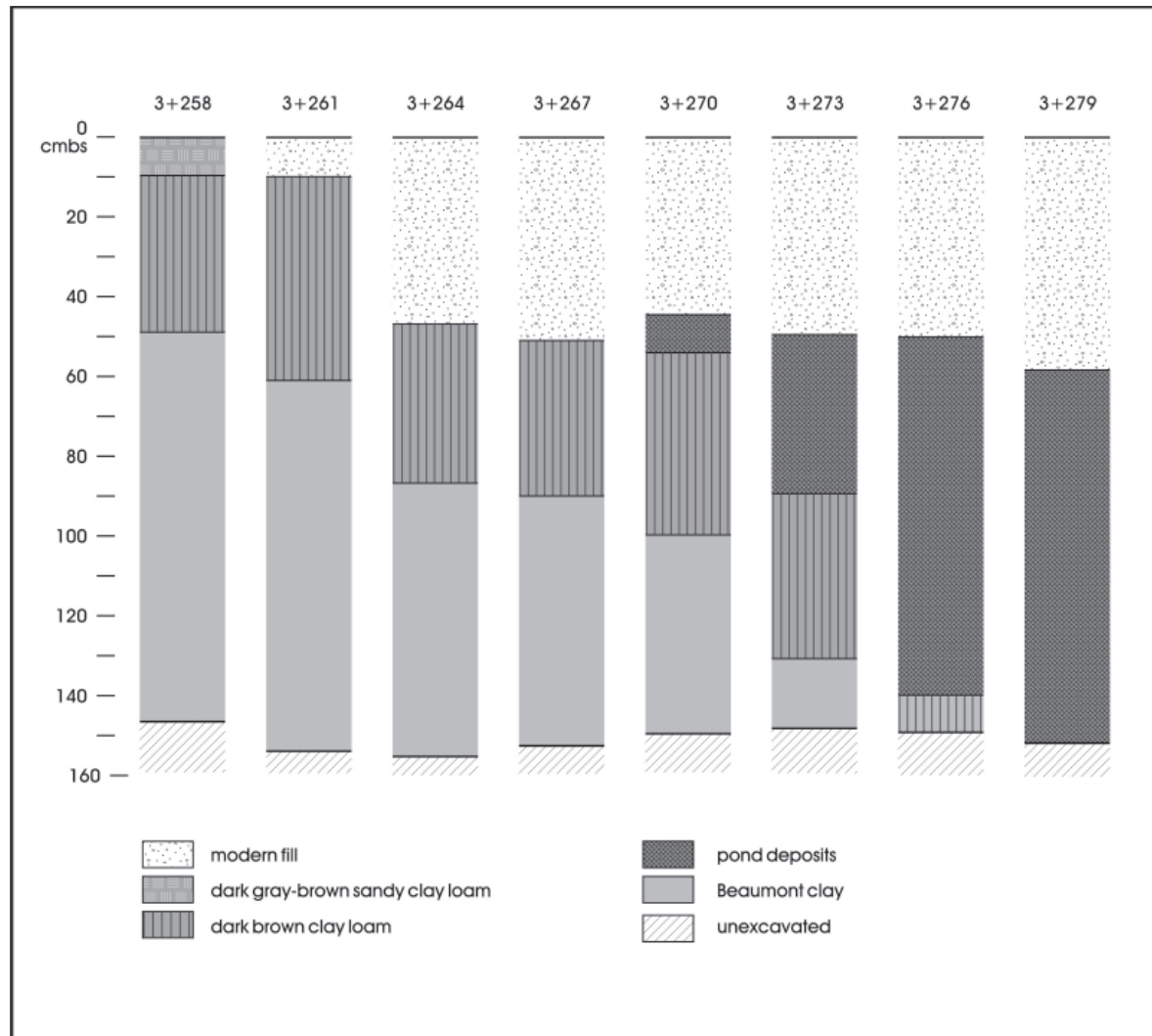


Figure 5-17. A series of eight 20-cm strip profiles of the north wall of Trench N, 3 m apart, from 3+258 to 3+279, documenting the western edge of the pond deposits and OZ K.

Trench E

Trench E was excavated during Phase II of the project. It began in an area that would later be Trench B at 34+152, and was dug east to a point just west of the edge of the stabilized bluff above the northern False Oso mudflats, then turned north to a point just south of Ocean Drive, roughly 155 m in total (Figure 5-1). Trench E was 120 cm wide and 180 cm deep. The first 20 m, on the western end of the trench, exposed undisturbed clay dune sediments, in which there were occasional shells. The remainder of the trench was dug through either disturbed sediments or modern fill including large amounts of modern trash, much of it in black plastic bags, in a sandy matrix which extended across most of the north-south portion of the trench. It appears that this garbage and sand mixture was used as landfill behind a stabilizing wall of jumbled, reused concrete slabs and blocks referred to as “riprap” by one of the TxDOT

inspectors (O. Hinze, personal communication 2004). This informal “seawall” is the method used in this area to protect the bluff from damage by storm surges on Corpus Christi Bay near the False Oso, and in some locations, along the bluff above the mudflats east of the site.

Trench F

Trench F was located in the parking lot of the restaurant north of Ocean Drive on a bluff above Corpus Christi Bay (Figure 5-1). This trench, dug to place a large storm sewer outlet, was 3 m wide and 2.95 m deep and extended 63 m from 18 m north of the Ocean Drive centerline at 3+328 to Corpus Christi Bay.

The complex layering of the North Dune was still clearly visible in Trench F. One dense concentration of sunray venus

clam and lightning whelk shells (Feature 18) was located in this trench (Figure 5-18). With the exception of a large sand-filled pit, roughly 8 m across and 2 m deep, beginning about 35 m north of the southern end of the trench, the dune deposits seen in this trench were undisturbed until about 72.3 m north of the Ocean Drive centerline. The pit had modern trash at the bottom, including a machine-made bottle. With the exception of Feature 18, no extensive prehistoric cultural deposits were noted.

Profiles of the trench showed that beneath roughly 30 cm of modern fill were clay dune deposits to the bottom of the trench. In a few places, the bottom layer seen in the trench showed the remarkable pale green color typical of sediments that have undergone gleying, that is, the sediments had been subjected to long periods of submersion in anaerobic conditions. This was confirmed when, after some time in the air, the sediment became a mottled tan and orange color, a phenomenon typical of gleyed sediments. In fact, Trench F broached the water table and the lowest part of the trench was submerged within a few minutes of finishing each segment (Figure 5-19). Figure 5-19 shows

the northern-most intact portion of the North Dune at approximately 71.6 m north of the Ocean Dr. centerline. Just north of this location the dune sediments were replaced with modern fill for another 5 m to the edge of the bluff above Corpus Christi Bay. Note the changes in the dune between Figure 5-18 and Figure 5-19, located about 52 m further north. Also note that the dark sediments are fading out as the dune reaches the original edge of the bluff above Corpus Christi Bay. No evidence of Beaumont clay was seen in any portion of this trench.

Trench H

Trench H was dug from Trench G northwest across the empty lot at the junction of Alameda (Figure 5-3). The portion of the trench nearest Trench G was expanded to make room for a large concrete manhole box. The upper 50 cm of the entire trench was highly disturbed. Beneath that was more of the dense black, highly organic clay deposits seen in Trenches D and G. Beneath these deposits, the Beaumont clay was very uneven, with the top varying from immediately beneath the disturbed sediments to as deep as 2 m bgs. No cultural deposits were seen in Trench H (Figure 5-20).

This trench crossed the part of the site where the Beaumont Ridge is highest. The Beaumont Clay Formation is almost completely impermeable to water so that, in the past, rain collected in the depressions on the undulating top of the ridge and remained there until it evaporated (Figure 5-20). These small ponds were usually very ephemeral except during rainy periods.

Trench I

Trench I was a 1-m-wide trench located roughly 50 cm east of Trench B, and originally extended from 33+530 to 34+145 (Figures 5-1 to 5-5). Unlike Trench B, this trench was only dug to about 120 cm bgs unless the Beaumont clay was encountered above this level. Not surprisingly, many of the larger features and all of the occupation zones seen in Trench B were also identified in Trench I. In addition, twenty-six features and two additional occupation zones were identified in this trench.

The stratigraphy of Trench I generally matched that of Trench B, except that matching paleosols of the north dune were somewhat lower in Trench I, as the layers of the entire dune slope to the east. On the other hand, ground surface in Trench I was consistently higher, sometimes as much as 50-60 cm higher, than the surface of the road through which Trench B was dug. This ground surface may be represented

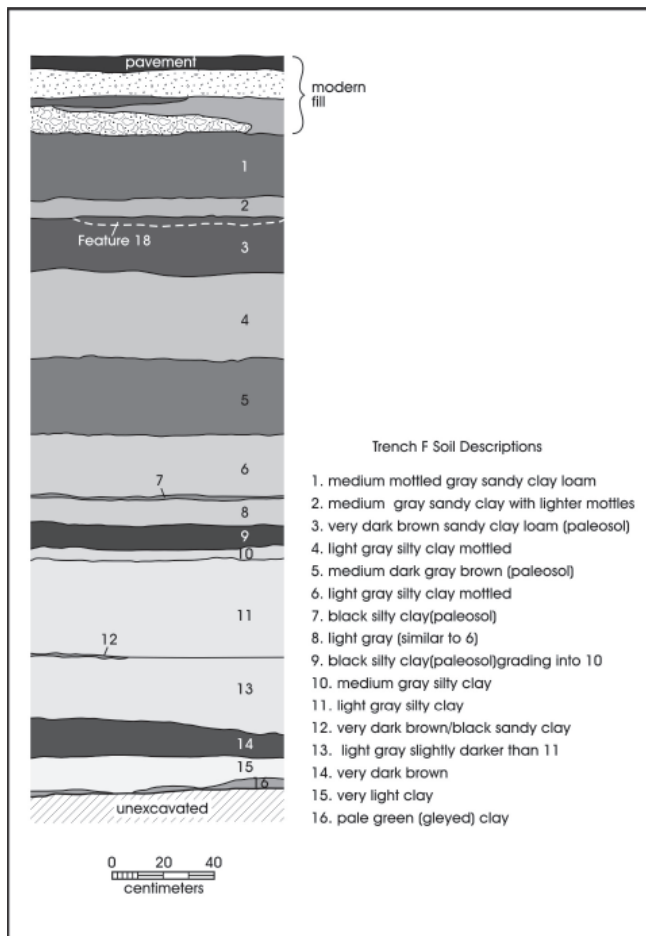


Figure 5-18. Profile of Trench F, at location of Feature 18.

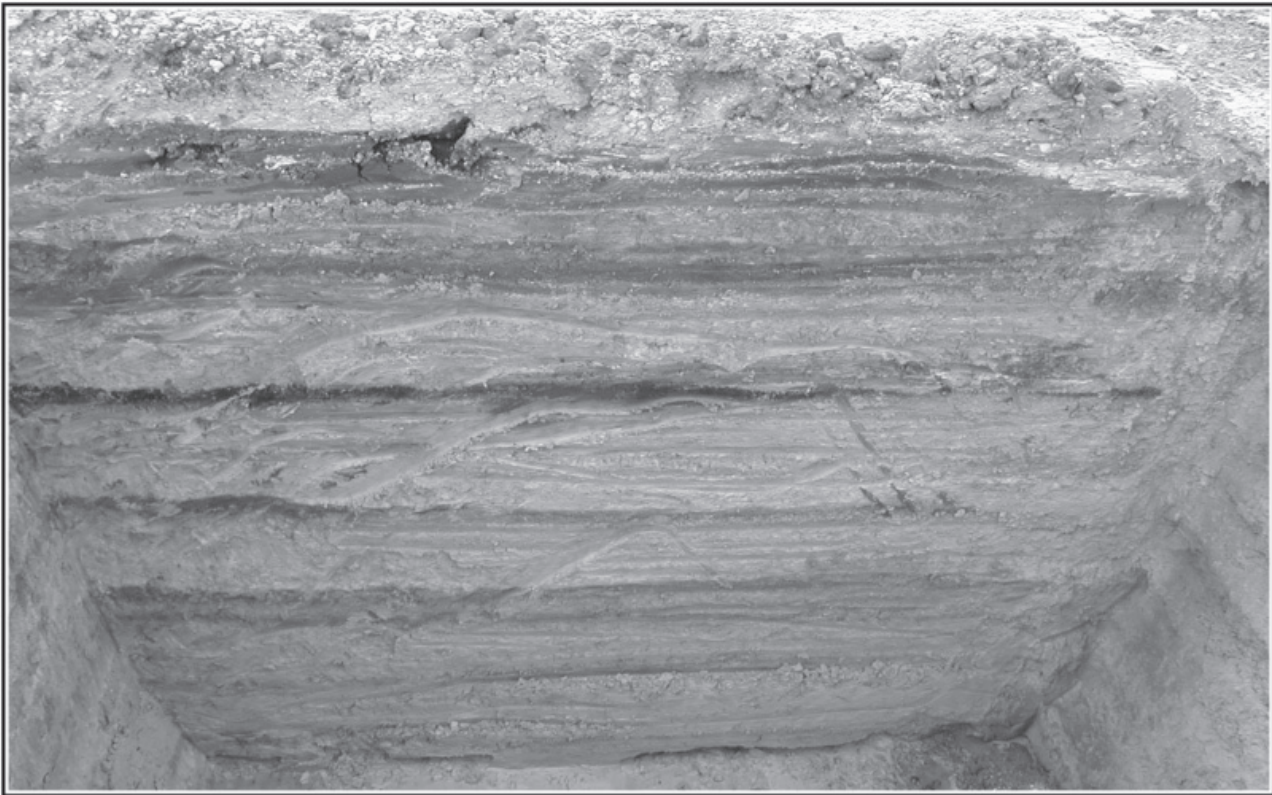


Figure 5-19. Photograph of northern-most point in the North Dune in Trench F, 71.6 m north of the Ocean Dr. centerline. Facing west.

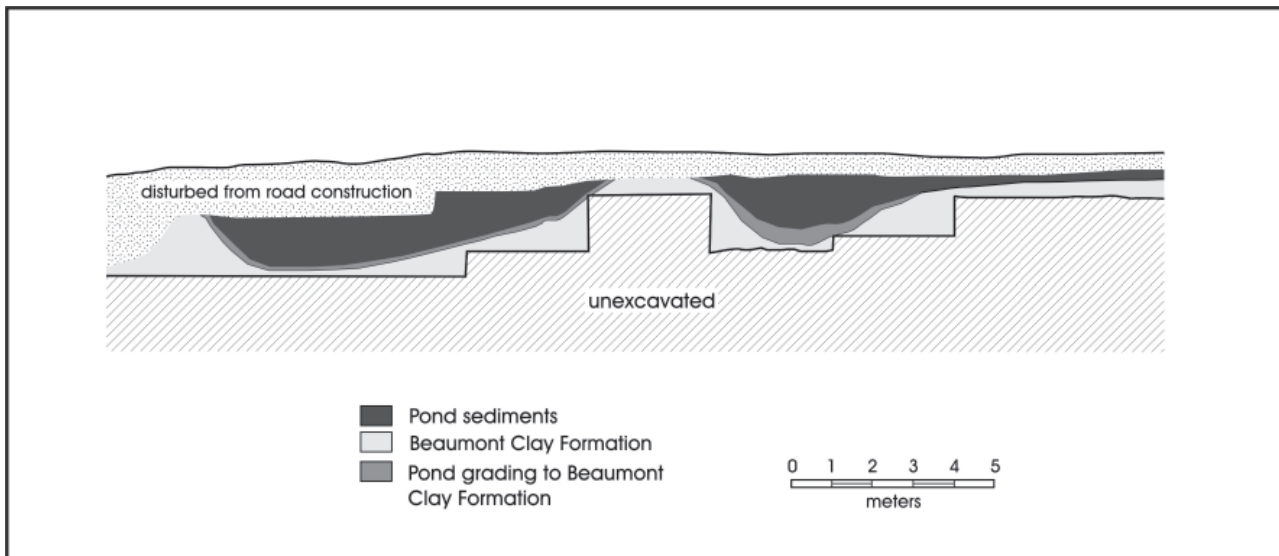


Figure 5-20. Profile of the north wall of Trench H, showing the uneven surface of the Beaumont Clay Formation and the black deposits left by small ephemeral ponds forming in lower elevations of the surface.

the natural ground surface in areas that had not been graded for landscaping. Thus, some features and occupation zones clearly visible in the upper part of Trench I were poorly defined or not seen in Trench B, presumably because they had been truncated by the road construction.

For example, OZ G was identified in Trench I but not in Trench B. Although the same 30 cm thick dark layer containing the occupation zone was seen in Trench B, it was noted at the time (during Phase I) as a dark layer with a thin scatter of shell fragments, charcoal and burned clay

specks. Since Trench B was not, at that time, continued to the north, it was not understood that this layer containing scattered cultural materials continued across more than 20 m of trench, and could be seen in the adjoining trench to the east. The cultural material in OZ G was also clearly far denser in Trench I.

Trench I was later continued from 34+145 on Ennis Joslin Road across Ocean Drive to a large manhole box at 3+361 on the northern edge of Ocean Drive. Although a number of large utility trenches had been dug in this area, undisturbed dune sediments were still present. Cultural materials were relatively rare within them, consisting for the most part of a few scattered shells, most of which were in the uppermost of the paleosols of the dune. While digging this trench, Burial 5, which consisted of a few disarticulated bone fragments in the backdirt of an old utility trench, was located at 34+151 (Figure 5-1). Burials 6 and 7, the disturbed but partially articulated remains of an adult and a juvenile, respectively, also were located in this trench at 34+212. More information about these burials can be found in a later section of this chapter.

Trench J

Trench J was a 50-cm-wide trench cut just east of Trench I to lay a temporary water line from about 33+530 to 33+618 (Figure 5-4). This trench was very shallow, usually no more than 70 cm bgs and for the most part did not encounter previously undisturbed sediments. No features or OZs were identified in this trench.

Trench K

Trench K was a footer trench for a retaining wall that began at the bridge abutment on the east side of Ennis Joslin Road at 33+353 and continued north to 33+478 (Figure 5-5). This trench was 2 m deep at the south end and became gradually shallower as it went north. The only undisturbed sediments seen were in the northern-most 5 m, at the very bottom of the trench. No prehistoric cultural deposits were noted.

Trench L

Trench L began at 33+544 on the east side of Ennis Joslin Road and extended west-southwest to the centerline, where it turned to the southwest, crossed Trenches A and C, and continued down the northern edge of Del Oso Drive to a large manhole box 38 m west of the centerline at 33+520 (Figure 5-4) where it turned south across Del Rio to a manhole box at 33+512. It was 1 m wide and varied in depth from 220 cm under Ennis Joslin to 120 cm near its western end. Five features (124, 125, 127, 128, and 135) and OZ H were identified in Trench L. A profile of the western end of the trench (Figure 5-21) shows that Features 128 and 135 are both sitting at the top of a light clay layer, while Feature 127 is within it. The profile becomes more complex as it moves east, suggesting that the eastern side is showing some beginnings of dune sediments. If we compare Figure 5-21 with Profile 2 in Figure 5-9, we can see that these profiles, only about 10 m apart, seem very different. It is important to note that these two profiles are in a transitional zone between the Beaumont Ridge and the South Dune. It is not possible to correlate a given stratum in

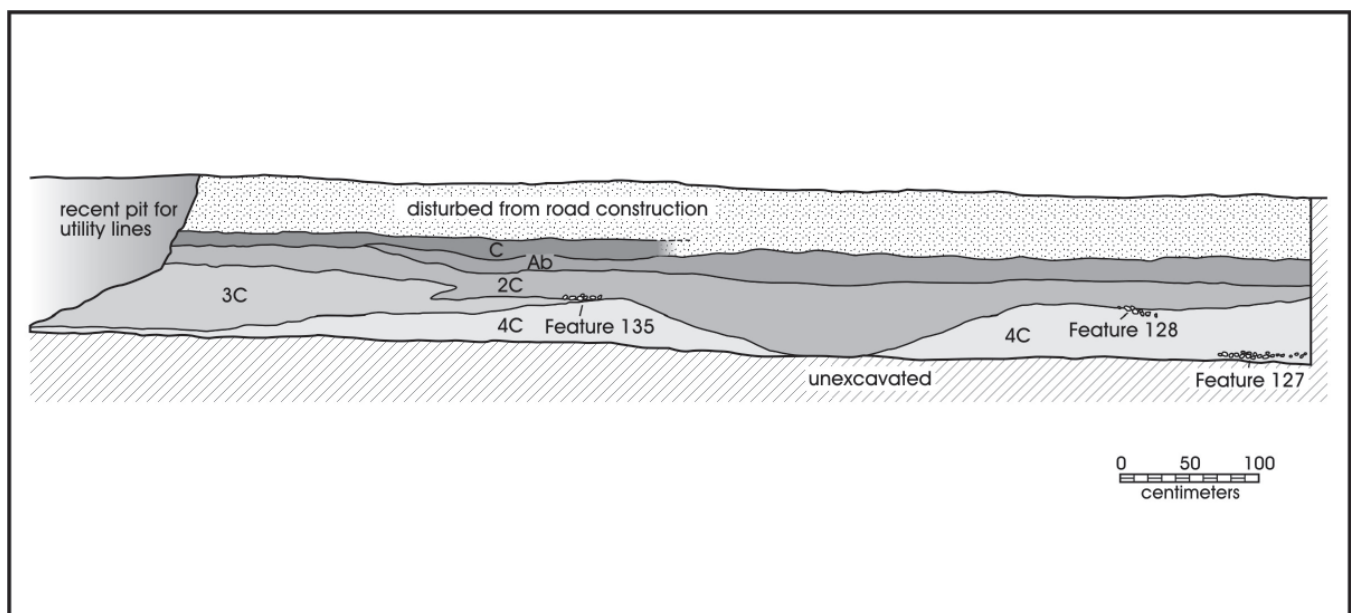


Figure 5-21. Profile of the south wall of the western end of Trench L at about 33+520 between 10 and 20 m west of the centerline, showing Features 127, 128, and 135.

Figure 5-21 to those seen in Profile 2 in Figure 5-9 with any certainty, given that it is not known how much sediment was removed above the intact deposits shown in Figure 5-21.

Trench M

Trench M crossed Ocean Drive near the intersection with Ennis Joslin Road (Figure 5-1), beginning and ending at storm sewer inlets. The trench extended from 3+295.5, on the northern curb of Ocean Drive, to a drainage inlet on the new curb line on the southern edge of the right turn lane from Ocean Drive onto Ennis Joslin Road (Figure 5-1). The trench was approximately 2 m wide and roughly 100-180 cm deep, with the deepest part of the trench located to the north. Most of Trench M transects an area where a large ephemeral pond is marked on old maps and is seen in old aerial photographs (e.g. 1934 aerial in Figure 5-8). The profile of the trench showed evidence of this pond, consisting of relatively uniform, dense, extremely black and rather glossy sediment that appeared to have a high organic content. When the removal of two old

storm sewer boxes was monitored not far east of Trench M (16.5 m west of the centerline at ca. 34+204 and 18 m west of the centerline at ca. 34+218, see Figure 5-2), this very black sediment was seen to extend to a depth of 2 m bgs in those areas. With the exception of one or two shells, no evidence of cultural activity was seen in this black sediment, however, the scattered remains of an American Bison calf (*Bison bison*), including several long bones with unsealed epiphyses, ribs, phalanges, and a portion of mandible with both permanent and deciduous teeth, were encountered between 90 and 110 cm bgs (Figure 5-22).

The bones were disarticulated, and scattered over an area of about 3 m north/south and an unknown distance east/west. There was no evidence of cultural activity associated with these bones, and since it was perceived as non-cultural it was not assigned a feature number. Nonetheless, the bones were collected for identification and closer examination in the laboratory. The results of the examination are presented in the faunal analysis section of Chapter 11.

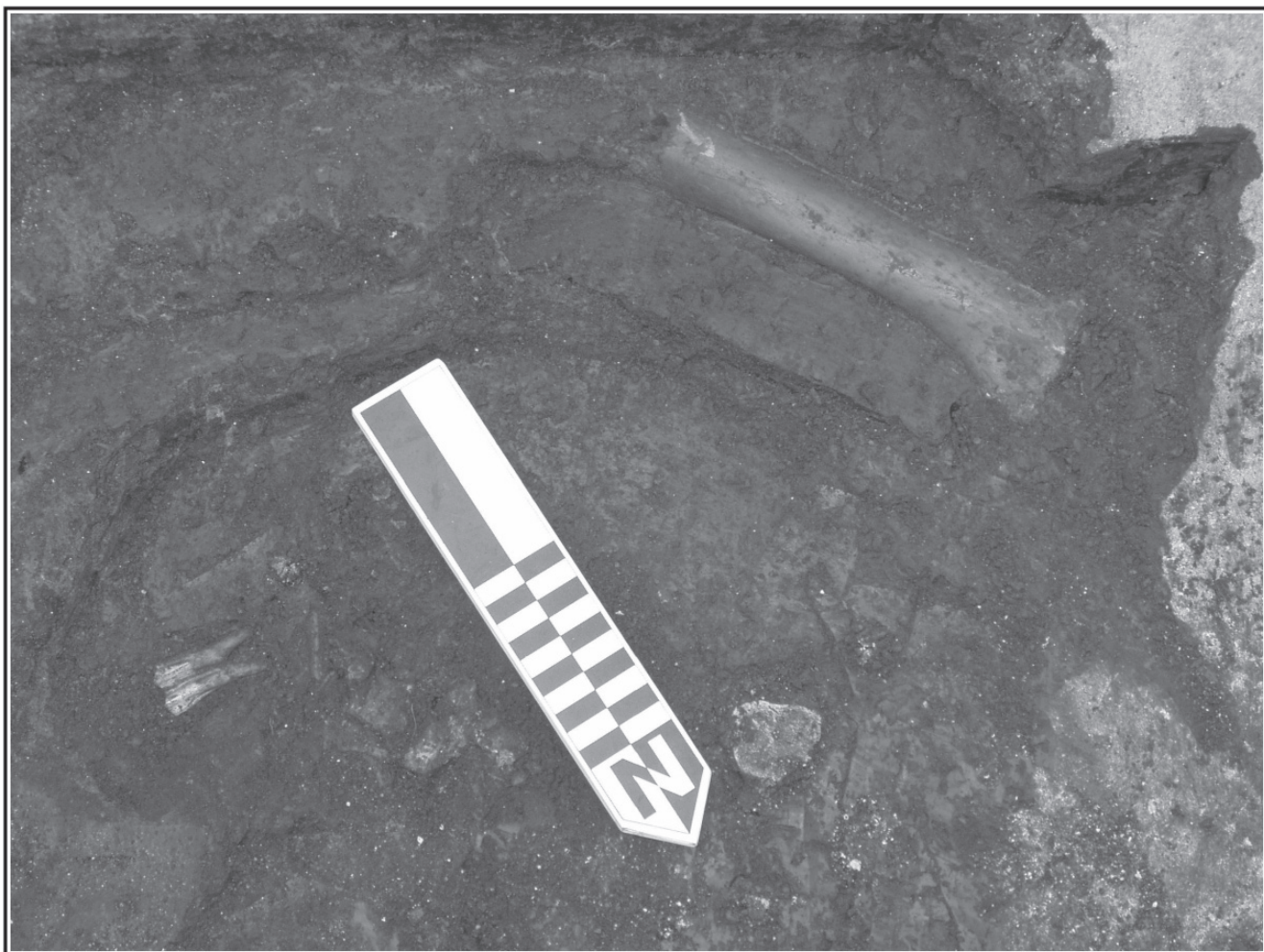


Figure 5-22. Immature bison bones in pond deposits in Trench M, ca. 1 m bgs.

Trench O

Trench O was dug to lay a temporary water line along the northern side of Ocean Drive (Figure 5-1). It ran from 3+215 to 3+400, approximately 60-100 cm north of Trench R. It was ca. 50 cm wide and 60 cm deep. The western 20 m of the trench was dug into Beaumont Clay, but with the exception of a small area between 3+271 and 3+280 all the sediments exposed were disturbed fill until ca. 3+345. East of this point, undisturbed sediments were seen in the bottom 10-15 cm of the trench. A single feature, Feature 163, was located in a previously undisturbed area at 3+275.

Trench P

Trench P was dug to replace the existing storm drain with a larger capacity pipe on the north side of Ocean Drive (Figure 5-1). It extended from 3+305 to 3+189. The trench was approximately 1.5 m wide and about 1.9 m deep. The trench extended about 20 cm below the original trench and roughly 30 cm wider on each side. The sediments observed in this trench consisted of dune deposit under a thick (40 to 60-cm) layer of fill and old road bed, until approximately 3+290, where Beaumont clay was observed in the bottom of the trench. By 3+280, the entire profile consisted of fill material above Beaumont clay, and the fill continued to the end of the trench at 3+189. Occasional shells were seen in undisturbed dune sediments, but no features were observed.

Trench Q

Trench Q was dug to connect a new sanitary sewer from the manhole at 3+360 to the restaurant north of the intersection of Ennis Joslin Road and Ocean Drive. It extended to a point about halfway across the parking lot, east of Trench F (Figure 5-1). It was ca. 90 cm wide and 120 cm deep. Two features (164 and 165) were identified in this trench. Both were just below the fill under the parking lot, in a paleosol that had been removed by road construction further south.

Trench R

Trench R was dug along the northern edge of Ocean Drive. It extended from 3+215 to 3+417.6 (Figure 5-1). It was ca. 2 m wide and varied in depth from 2.2 m bgs on the eastern end to ca. 0.9 m bgs on the western end. Trench R ran immediately to the north of Trench P and south of Trench O. The eastern end of this trench was located on a causeway which extends from the edge of the top of the dune deposit to the point where Ocean Drive crosses the False Oso. At this end of the trench there was approximately 70 cm of fill above an old roadbed (see Figure 5-7). Beneath this was another 80 cm of fill above

undisturbed dune deposit. The surface of these undisturbed deposits rose sharply as the trench continued west, towards the top of the dune. The lower roadbed disappeared and the fill became shallower until about 3+380, where the fill was only about 50 cm deep. West of this area, the upper surface of undisturbed dune deposit remained approximately the same depth to about 3+300, when it began to deepen, and by about 3+280, fill was present to the bottom of the trench. After about 3+290, the Beaumont Clay formation was seen below the fill in the trench. From this point to the end of the trench at 3+215, the profile consisted of fill material over Beaumont Clay. There were three features found in this trench: Features 166, 167, and 169. All three were at the surface of the undisturbed clay dune and had been partially disturbed by previous construction. A small bird bone bead was found in Feature 169 (see Artifacts section, below). Additionally, at the contact between disturbed and undisturbed sediments at 3+328, a single very small (ca. 1 cm) sherd of sandy-paste ceramic was found, apparently within the undisturbed context). Though there were a few shells near the ceramic, it was not associated with a distinct feature. This was the only piece of Native American ceramics recognized in the field during the project, although Martin (1930) describes Native ceramic fragments scattered on the surface in some areas of the site. Another piece of Native ceramics was located during floating of soil samples in the lab (see Artifact section).

Trench S

Trench S was dug to place a storm sewer outlet from two inlets on Del Oso to a shallow ditch leading to the unnamed drainage at the southern end of the site (Figure 5-5). It began at 33+519 at a storm sewer inlet, crossed under Del Oso to another inlet box and then turned to the south roughly parallel to Ennis Joslin Road. It was 1 m wide and varied in depth from 190 cm on the north end to ca. 30 cm at the southern end, where the storm drain would emerge from the ground. Beginning at about 33+480, the amount of fill above undisturbed clay dune deposit, which had been only about 20 cm deep, began to get deeper, as the surface of the undisturbed dune began to drop towards the unnamed drainage to the south. By 33+465, the trench was dug entirely through this fill. Feature 168 was located in this trench.

Trench T

Trench T extended north from a manhole at the end of Trench R (3+417.8) under a curb near the eastern end of the restaurant parking lot, extending 25.5 m and then turning to the east to connect to the water line for the Oso Pier building east of the restaurant (Figure 5-1). The trench crossed a concrete seawall at approximately 3+430 and extended another 15 m. The trench was 1 m wide to a point about 1 m north of the

curb and 50 cm wide beyond that point. It was 1.2 m deep. The trench exposed about 60 cm of parking lot and gravel over undisturbed clay dune. The surface of the clay dune dropped sharply as the trench extended east, and the depth of the fill increased. By about 3+425 there were no undisturbed sediments in the trench.

At the point immediately beneath the curve of the curb (see Figure 5-1) the backhoe operator dug out a large scoop to make his cut below the curb without breaking it, instead of scraping a few centimeters at a time as was his normal practice. Immediately beneath the curb he removed a large number of bovid bones. After removing this bone from the backdirt, the remaining bone within the trench was cleared enough to see that the bone bed (Feature 170) consisted of the articulated lower legs of several bovinds. Feature 170 will be discussed in more detail in the Features section of Chapter 6.

Trench U

Trench U was dug to place a sanitary sewer line between 34+002 and 34+053 on the west side of Ennis Joslin, west of Trench G (Figure 5-2). It was 1 m wide and ca. 1.2 m deep. The upper 65 cm of the trench was topsoil fill brought in for landscaping the lawn. Below this, the northern clay dune was clearly present and Features 141 and 142 were identified.

Sediment Cores

In August 2007, the Project Archeologist and a Field Technician accompanied the Project Geomorphologist to Corpus Christi to extract three sediment cores to be sub-sampled for a series of specialized analyses outlined in the project research design. This visit was the second attempt at the extraction of the cores after the first one, in June 2007, was not successful following catastrophic equipment failure.

Cores were to be extracted from the pond deposit, the north dune, and the south dune, respectively. Core segments were to be initially inspected in the field after extraction to ensure that they sampled intact deposits. Core 2 was extracted from a former pond deposit pictured in a 1934 aerial of the area (Figure 5-8). The north dune is represented by Core 4 that was collected from near the intersection of Ocean Drive and Ennis Joslin Road. The south dune, represented by Core 6, was cored along the south side of Ennis Joslin Road near its intersection with Del Oso Road. The core extraction locations are shown on Figures 5-1, 5-2 and 5-4.

The cores were collected from the project right-of-way immediately adjacent to Ennis Joslin Road, and not from

the summit of each dune. The core extraction process, the sub-sampling methodology and stratigraphic descriptions are presented in Appendix F. The results of the specialized analyses performed on the sediment samples are presented in Appendixes A-G.

Human Remains

Ricklis (1997) discovered one human burial during his project. We have designated that burial as Burial 1. During the CAR monitoring project, six additional burials were discovered: one complete, intact burial (Burial 2, Figure 5-2), three collections of disturbed, disarticulated human skeletal remains (Burials 3, 4, and 5; Figures 5-1 and 5-2), and two collections of partially disturbed human remains (Burials 6 and 7; Figure 5-1). Below, we briefly summarize the discovery and excavation of these remains. The analysis and a complete description of these remains were completed by CAR under a separate report in compliance with the NAGPRA inventory of human remains from the site. This report was submitted to the TxDOT-ENV as part of a separate deliverable of this project (Mahoney and Shafer 2003).

Burial 2

Richard B. Mahoney

Burial 2, the remains of a single individual, was encountered and excavated in December 2003. It was designated as Feature 61 (Figure 5-2). The feature consisted of the remains of adult individual in a tightly flexed position in a well-defined burial pit. The top of the skull was located 10.63 m east of the centerline at 34+53.9. The top of the skull was only 10 cm below the surface of previously undisturbed sediments. The burial was on its right side, with left hand near the face and right arm extended under the drawn-up knees.

During excavation, it was apparent that all skeletal elements were present, and this burial represents a single, primary interment. Taphonomic forces, however, have severely degraded the bone structure. Consequently, each individual element fractured or disintegrated upon removal. This fragmentation has precluded the observation and recordation of many metric and morphological traits.

Burial 3

On August 8, 2004, during monitoring of the removal of old utility lines in the vicinity of 34+038 about 16 m east of the centerline, several human bones were noted in the backfill of the trench (Figure 5-2). These consisted initially of a cranial fragment, a molar, and a possible mandibular fragment. All

bone recovered was disarticulated and fragmented and came from backdirt removed from a trench that was entirely within previous excavations to bury gas, water, and sewer lines.

Following instructions from TxDOT Environmental Affairs Division, the backdirt in that vicinity was screened with a ¼-inch mesh. In addition, the walls of the backhoe trench were examined for signs of previously undisturbed sediments and possible undisturbed portions of the burial represented by the recovered disarticulated bones. Both of the walls and the floor exhibited backfill sediments and there was no evidence of any intact soils.

Burial 4

On August 31, 2004 more human remains were discovered in backdirt from an old utility trench (Figure 5-1). The human remains were initially seen in an area at about 34+187, less than 1 m east of the centerline, when a layer of fill was knocked into a previously dug trench. Bones were initially identified as human from a distal humerus, a proximal radius, and a patella, all of which were readily identified in the field. Although all bone was disarticulated and fragmented, they were collectively designated Burial 4. It was later discovered that the bone all came from a trench (Feature 136) that had been machine-dug at some time in the past into approximately 120 cm of builders' sand that had been filled over an old storm sewer that runs roughly east-west across the traffic islands located just south of Ocean Drive. Feature 136 was later completely excavated and its contents screened. It extended between 34+185.5 and 34+187.7, from 0 m to 1 m west of the centerline.

All bone was disarticulated and fragmented and was found in a gray sandy clay matrix that was very different from any of the sediments, whether disturbed or undisturbed, seen in the vicinity. This matrix had clearly been disturbed and contained small limestone rocks and gravel (which is not locally available rock), as well as metal and plastic fragments. Within this matrix, were clumps of dense clay that more closely resembled the sediments usually seen at the site.

The evidence suggests that Burial 4 was encountered by a construction crew at some time after the storm sewer was constructed. In order to rebury these remains, the small trench was dug into the sand above the storm sewer, and the remains, along with a sandy matrix that did not come from the immediate area, were deposited and covered with several layers of fill. The original location of Burial 4 cannot be ascertained. Although most of the bones were fragmented, they were in relatively good condition.

Burial 5

Burial 5, identified on September 9, 2004, consisted of a few highly fragmented human bones found in the fill above an old cement sanitary sewer line. The first bone was found 16 m east of the centerline at 34+152 (Figure 5-1). The matrix in the trench in the immediate vicinity of the location of the first bone was screened and several other human bones were collected. The bone was found in a sandy matrix immediately above the sewer pipe, with a second layer of sandy matrix above it. The human bones were only found in the matrix immediately above the pipe, in an area extending north about 70 cm from the location of the original find. Backdirt in an area about 2 m on each side of this location was closely examined, as instructed by TxDOT Environmental Affairs Division, but no other human bones were recovered at that time. However, several days later, the excavation around the old sewer trench was expanded and several more bones that appeared to be human were identified at 34+155. Again, the matrix from the trench in the immediate area was screened and the back dirt in an area about 2 m to the north of this area and south to the portion of the trench already examined (roughly 30-40 cm), was closely examined. No other possible human remains were recovered.

Burial 5 is clearly another case of previous construction activities impacting on a burial. Unlike the bones of Burial 4, these bones were in poor condition, some showing indications of atmospheric weathering, suggesting that the bones had, at some time in the past, been exposed above the ground surface for some time. The bones were found in two locations approximately 3 m apart, but the general appearance, color, and condition, as well as the fact that they were found in the same layer of fill material over the sewer pipe, suggested they were from a single burial. Since the bone was scattered over a wide area, and as the fragments were quite small, it is possible that the construction workers were not aware that they had disturbed a human burial.

Burials 6 and 7

Russell D. Greaves

On October 13, 2004, during expansion of Trench I around a broken segment of the same concrete sewer line along which the disturbed bones of Burial 5 were located, the partially disturbed burials of two individuals were encountered. These burials were 29.1-29.3 m east of the centerline at 34+208.5 to 34+209.3 (Figure 5-1). Initially one long bone fragment was encountered at the margin of the trench. On the basis of the size and morphology, it was determined that the disturbed bone was probably human and CAR notified TxDOT that another probable burial had been found.

Although the first bone identified seemed to be that of an adult, additional careful troweling uncovered very small ribs that appeared to be in anatomical position. A deciduous tooth root and what was later identified as an unerupted permanent premolar were found approximately where the cranium might have been. Additional excavation was directed toward uncovering the horizontal extent of this burial. The partially disturbed burial of a small child was uncovered. The long bones and ribs were, as far as could be determined, articulated, however, the bones were in extremely fragile condition, and the cranium, with the exception of the unerupted tooth cap and the deciduous tooth root, was not evident.

As the area around the burial was excavated, additional ribs, pelvis, and long bones were located just eastward of the child burial and at a slightly higher elevation. These bones were clearly those of an adult. They were not in anatomical position, and it was initially unclear what form of commingling was represented between the juvenile and adult remains. At the end of the day the remains were photographed and mapped.

The following day, heavy rain prevented any attempts at excavation of the burials. Excavation was resumed on October 15. Mechanical removal of some of the overburden was performed to determine whether additional remains could be identified or if a grave pit was visible within the soil. Backhoe excavation was stopped well above the elevation where the highest bone had been encountered. Recent disturbances seen along the northwest, southwest and southern sides were very close to the remaining bones. It is likely that several additional elements on the western side were removed during one or more previous construction efforts. No grave pit outline was identified.

The child burial was at a lower elevation (74-82 centimeters below surface [cmbs]) than the adult remains (primarily 62-70 cmbs, with one fragment recovered at 82 cmbs). As excavation continued, it appeared that the child's burial had been excavated through part of the adult grave. Even the adult bones that were in the intact prehistoric sediments were jumbled. The disarticulation of the adult bones could represent a secondary burial, but in any case it appears that the adult remains had been disturbed by digging of the child's grave, and then again by modern construction activities. The adult burial was designated Burial 6 and the juvenile as Burial 7.

In summary, Burials 6 and 7 appear to be the disturbed remains of burials which were not contemporaneous. The adult burial may have been secondary, but this cannot be

clearly established, due to the highly disturbed nature of the context and evidence of considerable previous construction activity nearby. Burial 7, though in very poor condition, may have been essentially undisturbed. Preliminary examination of the unerupted tooth cap recovered from Burial 7 suggests that it belonged to a child between the ages of 3.5 and 6.5 years.

It is likely that other bones from Burial 6 were removed during that previous construction. It is even possible that the bones found roughly 50 m south of Burial 6, which were designated Burial 5, are some of those missing from Burial 6. It is highly unlikely that it will be possible to prove or disprove this speculation, and only the fact that Burial 6 was apparently disturbed by placement of the same sewer line near which Burial 5 was found makes the suggestion tenable.

Artifacts

The methods used during the monitoring project at 41NU2, combined with the nature of most of the artifacts, made recognition difficult. Most artifacts that were recognized are made of shell, and most were surrounded by hundreds, if not thousands of shells that had not been modified. There is no doubt that many artifacts were present but not seen from the top of the trench. The small number of artifacts presented here should not be considered evidence that few artifacts were, or are, present at 41NU2. This section can serve to show the sorts of artifacts that can be expected at similar sites, but it cannot serve to suggest what proportion of each type is likely to be found in the future.

A total of 123 artifacts was identified during this project, of which 33 were lithics (26.8 percent), 7 were bone (5.7 percent) and 83 were shell (67.5 percent). These artifacts will be described and discussed below.

Lithic Artifacts

Jason B. Perez

Thirty-three lithic artifacts were recovered from 41NU2 (Table 5-1). Of these, 25 (76%) came from feature matrix and were recovered during flotation. Eleven features (1, 41, 57, 65, 73, 75, 81, 90, 110, 117, 166) contained lithic artifacts. Of these, Feature 65 had ten specimens, Feature 73 had three, Features 57, 75 and 90 had two each and all others had one specimen each. All of the feature associated lithic artifacts are unmodified lithic debitage. Of the 33 lithic artifacts, only two are dart points, all others are unmodified lithic debitage.

Table 5-1. Lithic Artifacts from 41NU2

FS # Lot	Feature	Depth (cm bgs)	Artifact Category	Completeness	Cortex	Flake Type	Max Dimension (mm)	Comments
129	near F-40	83	dart point	proximal frag	–	–	–	Possible Matamoros dart point; heavy rejuvenation; stem appears to have a slight neck which does not fit with Matamoros type
406	–	backdirt	dart point	complete	–	–	–	Matamoros or heavily rejuvenated Tortugas dart point
238-1	57	60-70	unmodified deb.	proximal frag	secondary	hard hammerstone	19.2	Coarse grained chert with corticate platform
238-2	57	60-70	unmodified deb.	complete	secondary	hard hammerstone	7.5	Small corticate platform flake
449	–	72	unmodified deb.	distal frag	tertiary	indeterminate	17.3	Fine-grained chert
270	65	45	unmodified deb.	complete	tertiary	pressure	4.3	Single faceted flake
534	166	50	unmodified deb.	longitudinal frag	tertiary	pressure	7.4	Dark fine-grained chert
447	–	72	unmodified deb.	proximal frag	secondary	hard hammerstone	16.9	Single faceted platform
130	–	83	unmodified deb.	proximal frag	tertiary	indeterminate	22.5	Two platform facets; fine-grained chert
229	81	80	unmodified deb.	angular debris	secondary	indeterminate	n/a	–
270	65	45	unmodified deb.	complete	tertiary	pressure	5.7	Multiple faceted
270	65	45	unmodified deb.	complete	tertiary	pressure	5.3	Single faceted platform
270	65	45	unmodified deb.	complete	tertiary	pressure	4.4	Single faceted platform
270	65	45	unmodified deb.	complete	tertiary	pressure	5.8	Single faceted platform
270	65	45	unmodified deb.	complete	tertiary	pressure	3.8	Single faceted platform
270	65	45	unmodified deb.	proximal frag	tertiary	pressure	4.4	Single faceted platform
270	65	45	unmodified deb.	distal frag	tertiary	pressure	4.3	–
270	65	45	unmodified deb.	distal frag	tertiary	pressure	5	–
246	90	90-110	unmodified deb.	longitudinal frag	tertiary	pressure	5	Longitudinally split
246	90	90-110	unmodified deb.	angular debris	tertiary	indeterminate	–	–
Lot 363	–	–	unmodified deb.	complete	tertiary	biface thinning	26.8	25 mm long multi-faceted
209	75	71-90	unmodified deb.	medial frag	secondary	indeterminate	13.8	Medial frag; fine-grained chert
209	75	71-90	unmodified deb.	medial frag	tertiary	indeterminate	4.3	Medial; fine-grained chert
321	117	32-42	unmodified deb.	complete flake	tertiary	pressure	6.3	Multiple faceted platform
7	1	95-100	unmodified deb.	medial frag	tertiary	indeterminate	7.7	Medial frag; fine-grained chert
450	–	73	unmodified deb.	medial frag	tertiary	indeterminate	12.4	Medial frag; fine-grained chert
199	73	45	unmodified deb.	proximal frag	tertiary	pressure	6.5	Proximal frag, single faceted
199	73	45	unmodified deb.	angular debris	tertiary	indeterminate	–	Small debris
199	73	45	unmodified deb.	complete	tertiary	pressure	9.8	Single faceted platform
135	41	74-78	unmodified deb.	proximal frag	tertiary	biface	15.6	Moderately large proximal frag;
269	65	60	unmodified deb.	distal frag	secondary	indeterminate	19.8	Moderately large distal flake frag
293	110	155	unmodified deb.	complete	secondary	indeterminate	16.7	Moderately large flake
Lot 42	–	–	unmodified deb.	medial frag	tertiary	pressure	6.5	–

Dart Points

Both dart points are complete and neither was associated with a feature. Of the two points, one, a triangular specimen, appears to be a Tortugas point made of yellow fine-grained chert (Figure 5-23a). While the specimens is relatively short (23.3 mm) and wide (base width 25 mm), both faces contain broad base thinning flake scars that are similar in width to those found on Tortugas rather than the smaller Matamoros points (see Tomka

2002). The margins of the blade have rather steep retouch although it is rather unsystematic and does not classify as beveling. The point has been heavily retouched although examination under both short and long-wave ultraviolet light does not show differential patina of the point. This indicates that the reworking occurred during the use-life of the specimens or shortly thereafter. This point was found on the surface in the North Dune area near 34+041, after removal of a pile of backdirt. Thus, no provenience information is available.

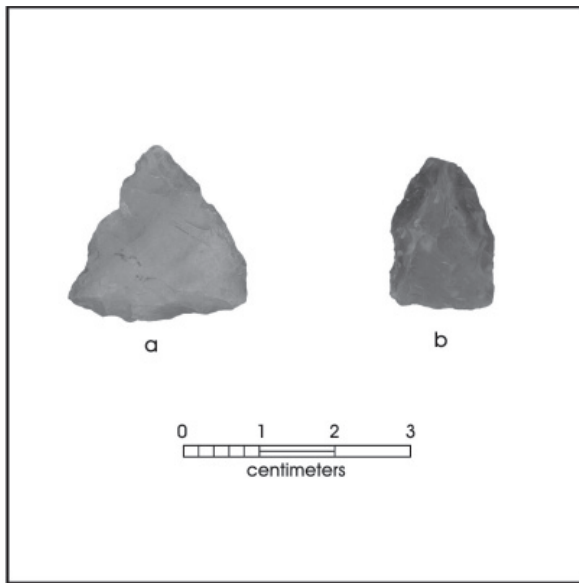


Figure 5-23 Heavily retouched dart points: a) Tortugas; b) untyped slightly expanding stemmed point.

The second specimen is also roughly triangular although its final appearance is the product of heavy rejuvenation. It is an untyped straight-based and parallel to very slightly expanding stemmed dart point made on translucent brownish flint (Figure 5-23b). Its blade has been so heavily reworked that it is actually shorter (10.4 mm) than its stem (10.9 mm). A small portion of its distal tip is missing due to an impact fracture. The slightly expanding stem (neck width 13.4 mm; base width 14.5 mm) gives it the appearance of a Darl point although this typological assignment cannot be made with certainty. It was recovered near Feature 40, in OZ-B. A charcoal sample from this OZ was dated to 2300-1880 B.P., which cross-cuts the Transitional Archaic, and seems consistent with the possibility that the point is a Darl.

Unmodified Debitage

Thirty-one pieces of unmodified lithic debitage have been recovered from the site. Only five of these are not associated with features (Table 5-1). Of the 31 pieces, three are angular debris, one of them has cortex the other two are decorticate. Tertiary unmodified debitage (n=25) greatly outnumber secondary pieces. On average they are small (mean maximum dimension = 8.7 mm) and have small platforms and diffused bulbs of percussion. These traits indicate that the majority of them were removed by pressure flaking either during tool rejuvenation or the final stages of tool manufacture. Secondary flakes are on average larger (15.6 mm) than tertiary specimens. The fact that two of them have corticated platforms in combination with their larger size suggests that the raw materials they represent may not have been obtained

far from the site, though to date no source for chert cobbles closer than the one described by Chandler (1984:27) has been found. Chandler's (1984) report described a chert gravel outcrop being eroded by the Nueces River at a point about 54 km upstream of the mouth of the river in Nueces Bay and about 57 km from 41NU2.

Bone Artifacts

Eight artifacts made of bone, most of them fragmentary, were recovered during the project. The complete list is provided in Table 5-2. Bone artifacts of particular interest were two beads, and two bone tools made on the metatarsals of artiodactyls.

Beads

Bone beads are inherently fragile and small, and therefore it is not surprising, that only one was identified during monitoring. The one complete bead was made out of bird bone, and was found in Feature 169 (Figure 5-24a). It is unusually short for a bird bone bead, but wear-polish on both sides indicate the bead was not broken.

A bone bead with unusually detailed workmanship was found in the soil sample from Feature 75 (Figure 5-24b). This bead was carved out of a solid piece of mammalian bone. The faces are concave, the hole is large and appears quite circular. The outer rim of the bead has a groove cut into it. The remaining bead fragment appears to be not quite half of the original bead, which would have had a diameter of roughly 7.8 mm.

Metatarsal tools

Two bone tools were made of the proximal diaphysis of the metatarsal of artiodactyls, probably deer (*Odocoileus virginianus*). The first of these was an isolated find, still intact until the weight of the Gradall bucket slid over it. This is an unusually well-made tool, symmetrically smoothed, highly polished, still very sharp (Figure 5-24c). Only minor chemically and biologically induced damage is visible on its surface. The specimen has an oval cross-section. Tools like these have traditionally been referred to as awls. Microscopic scratches running parallel to the long axis are present across 15.5 mm of the distal end. These appear to have been acquired during use.

The second specimen is a fragmentary metatarsal tool similar to the one described above although it may have been somewhat shorter when complete (Figure 5-24d). It is encrusted with a layer of calcium carbonate and the

Table 5-2. Bone Artifacts from 41NU2

Provenience	Feature	Length (mm)	Width (mm)	Wgt. (g)	Notes
15 m N of CL at 409.1, 140 cmbs	169	5.9	6.0	0.1	Small, complete bird bone bead. Figure 5-25a;
13.2 E of CL at 34+090.5, 71-90 cmbs	75	6.5	2.8	0.3	Bone bead fragment, not made from bird bone, but the fragment is too small to be identifiable. The bead was very well made, with concave faces, and a groove cut on the outer edge. It would have been ca. 7 mm in diameter. Figure 5-25b;
11.09 m E of CL at 33+595.2, 64 cmbs	–	116.2	15.9	11.9	Beautifully made bone tool, complete until it was broken by the Gradall. Highly polished and still with an extremely sharp point that was broken from the tip. Figure 5-25c;
14.6 m E CL at 34+095.2, 80 cmbs	74	65.4	12.5	4.1	Bone tool made on a fragment of artiodactyl metatarsal, with small areas of the anterior left on the tool. Well smoothed, with distal end broken. Figure 5-25d;
1 m W of CL at 34+185.51. 20-120 cmbs	136	54.0	6.4	1.2	Highly polished fish bone, possibly a spine, smoothed and shaped, found in the same backdirt in which the disarticulated elements of Burial 4 were located. Figure 5-25e;
14.6 m E of CL at 34+139.2, 25-32 cmbs	114	4.5	9.0	0.1	Fragment of a bone tool, heavily burned and polished;
10.73 m E of CL at 34+012, 84 cmbs	–	27.6	11.7	1.3	Small fragment of a bone tool, with clear shaping and smoothing that is too small to identify further;

bone surface appears more weathered than the complete specimen. It may not have been finished, as remnants of the anterior portion of the metatarsal are still present on the bone and the polish is uneven.

Other Bone Artifacts

One artifact of unknown use was made on a fish bone, probably a dorsal or haemal spine of a large fish. The bone is very highly polished and lightly burned (Figure 5-24e). It appears to have been only minimally shaped. The bone is broken on both ends, but a series of microscopic scratches cutting through the polish and running perpendicular to the length of the bone can be seen at the distal end near the break. This tool was found in the same backdirt which contained Burial 4. It is impossible to know if it was associated with the burial.

One fragment of a bone tool only 11.5 mm long was recovered from the soil sample from Feature 114. The fragment is heavily burned and highly polished, and appears to be a segment near the distal end of a pointed tool. Another fragment from Trench B in the North Dune shows clear evidence of shaping and smoothing but is too small to identify further.

Shell

There were 83 modified shells identified during this project, although as mentioned above, these artifacts were often surrounded by many shells that were simply the remains of a meal. Six species of mollusk showed evidence of tool manufacture or other deliberate modification. By far the most common shells modified into tools were those of Sunray Venus clams (*Macrocallista nimbosa*, n=38) and Lightning Whelks (*Busycon sinistrum*, n=32). Nine perforated Eastern Oyster shells (*Crassostrea virginica*) and one oyster shell ground into a tool were also recovered. One shell each of Florida Horse Conch (*Pleuroploca gigantea*), Southern Quahog (*Mercenaria campechiensis*) and Banded Tulip (*Fasciolaria lillium*) also showed modifications. All modified shells are listed in Table 5-3. In addition, some modified shells of particular interest are shown in Figures 5-25 through 5-32 and discussed below. The discussion of the modified shells is organized by shell species rather than tool form.

Modified Sunray Venus shells

The thin but dense shells of Sunray Venus clams were commonly used for tools along the Texas coast, in many cases replacing stone tools where chert and other stone was so difficult to obtain (Ricklis 2004). The shells of

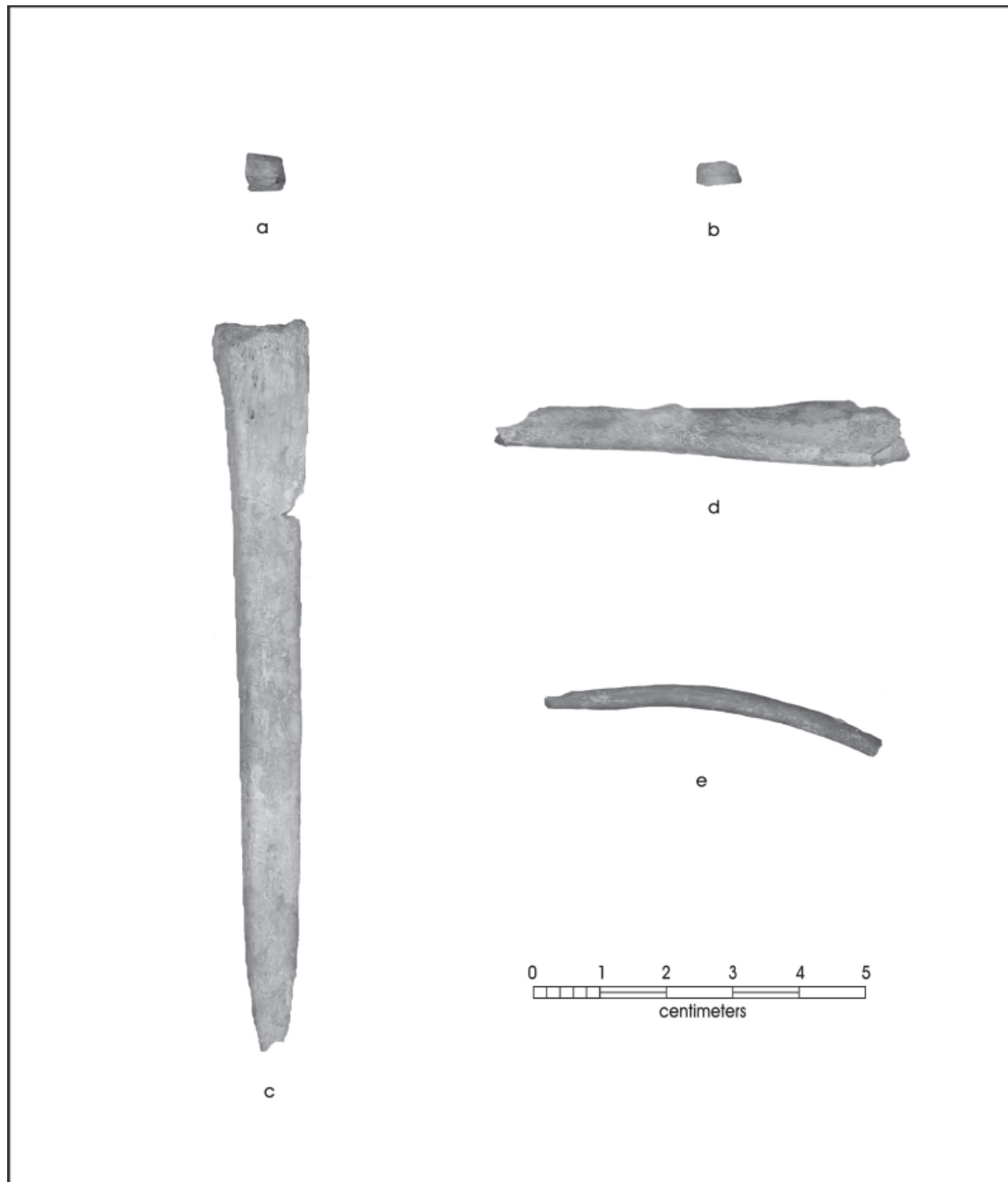


Figure 5-24. Bone artifacts: a) bird bone bead from Feature 169; b) bone bead fragment from Feature 75; c) complete metatarsal tool; d) fragmentary metatarsal tool; e) fish bone tool.

Sunrays can be ground or chipped to shape them into useful impediments.

Ground shell tools. Eighteen ground Sunray shells were identified during the project (Table 5-3), six of the more complete ones are discussed in some detail. Tools made by grinding Sunray shells typically made use of the way that Sunray shells naturally break, which often leaves a long edge

with some of the harder outer shell removed. In these cases the inner shell was often ground down so the edge became very sharp. Figure 5-25a shows a tool in which an S-curved break was sharpened by grinding. A large portion of this edge has broken off, but some of the grinding is still visible under the microscope. Also visible are a number of deep scratches running perpendicular to the long axis of the shell, suggesting that after the sharp edge was broken away, the tool was used at least briefly as a scraper.

Another tool, roughly triangular in shape was also ground (Figure 5-25b), though again, most of the ground edge has broken away. Both the grinding and the scratches that are presumably part of the use-wear on this tool are readily visible under the microscope. The third specimen (Figure 5-25c) is a rectangular shell fragment that has been ground on two edges. The third edge is the naturally thinned ventral edge of the shell, and the fourth edge is a break. The fourth ground shell tool (Figure 5-25d) is a triangular shaped tool, with the hinge plate forming one edge. The single ground edge is still sharp enough to break the skin and shows little use damage, unlike the tools described above.

The fifth specimen (Figure 5-25e) is a ground Sunray shell tool of particular interest because it was found among the cattle bone in Feature 170. It is a long sliver of Sunray shell ground along one edge. Although most of the evidence of grinding has been removed by post-depositional breakage, grinding is clearly visible on the distal end. The development of a fairly thick patina prevents examination for use-wear scratches. Because the cattle bone were dumped into an erosional gully and later covered by a natural infilling process, this shell tool

and one other found in the same location (see below) were not likely associated with the cattle bone discard. However, their presence does open the possibility that Feature 170 represents a Protohistoric component at the site.

The sixth tool (Figure 5-25f) was recovered from Feature 151. It was made by breaking away the shell hinge plate and then grinding the hinge plate into a long thin tool. The end has broken off so there is no way to tell if it was pointed or flat.

Chipped Shell Tools. The shells of Sunrays are very dense, though quite thin, and the Native Americans of the Texas Coast found that they could be chipped in much the same way as chert. Ten chipped Sunray shell tools were identified. Four are discussed in some detail below.

Most chipped shell tools seem to be made by simply chipping the ventral edges of the shells and leaving the rest or most of the rest of the shell intact. The specimens shown in Figures 5-26a and 5-26b are broken examples of chipped shell tools with unifacial chipping along the ventral edges of the shells.

Table 5-3. Modified Shell Artifacts from 41NU2

Provenience	Feature	Length (mm)	Width (mm)	Wgt. (g)	Notes
Sunray Venus Clam (<i>Macrocallista nimbosa</i>)					
Ground					
Backdirt near 34+062 on west side	–	95.2	42.6	24.2	A large Sunray with a curved break that has been ground to an edge. Most of this edge has been broken away but the remnants of the grinding is still visible. Figure 5-26a;
3.28 m E of CL at 33+919.2	76	76.9	37.5	7.1	Large triangular fragment of Sunray shell, ground to an edge on all sides, with two sides showing some random chipping that may be use wear. Figure 5-26b;
14.6 m E of CL at 34.074.1, 49 cmbs	65	37.8	29.9	4.6	Sunray shell fragment, roughly rectangular though it is broken on two edges. Two edges have been ground smooth. Figure 5-26c;
14.3 m E of CL at 34+047.7, 46 cmbs	98	57.8	30.3	12.5	Sunray Venus shell snapped into a triangle, with one edge sharpened by grinding. Figure 5-26d;
19.7 m N of CL at 3+412.6, 97 cmbs	170	52.4	10.2	2.9	Long Sunray fragment ground along one edge. Only the distal end still shows grinding. Found in association with cattle bone in Ftr. 170, but could easily have been washed into the gully. Figure 5-26e;
12.2 W of CL at 34+058.5, 57-62	151	43.7	14.8	–	Sunray shell, with hinge plate extending from the umbo ground into a long thin extension, which has broken. Figure 5-26f;
11.3 m E of CL at 34+084.5, 45 cmbs	73	22.3	16.5	1.2	Sunray shell tool made on a small triangular chip, with ground edges;
13.2 m E of CL at 34+090.5, 71-90 cmbs	75	65.2	24.5	14.8	A Sunray shell snapped parallel to the umbo, then ground to sharpen;
14.7 m E of CL at 33+589, 110 cmbs	84	26.5	15.7	2.0	Sunray shell fragment, roughly rectangular though it is broken on two edges. Two edges have been chipped to make an edge;
14.7 m E of CL at 34+060.4, 42 cmbs	100	20.8	21.5	1.9	A Sunray shell fragment with one edge ground to sharpen;
14.7 m E of 34+064, 42 cmbs	100	16.0	10.0	0.2	Thin fragment of a Sunray shell ground to a very sharp edge on two sides;

Table 5-3. Continued...

Provenience	Feature	Length (mm)	Width (mm)	Wgt. (g)	Notes
Sunray Venus Clam (<i>Macrocallista nimbosa</i>)					
Ground					
14.5 m E of CL at 34+136	114	23.4	13.1	1.0	Small fragment of Sunray shell that appears to have been sharpened, but also is too weathered to tell for sure;
14.6 m E of CL at 34+139.2, 25-32 cmbs	114	21.6	15.0	1.2	Sunray shell fragment, roughly rectangular though it is broken on three edges, and ground to an edge. The entire tool fragment is heavily patinated;
14.4 m E of the CL at 33+902.6, 110-115 cmbs	121	22.0	16.3	1.3	Fragment of Sunray shell ground on one edge to sharpen. This appears to be a fragment of a larger tool;
2 m E of Cl at 34+174.5, 55-60 cmbs	139	16.2	13.9	0.3	Small chip of Sunray shell, ground into a roughly square shape. Two edges show microscopic chipping that may be use wear;
2 m E of Cl at 34+174.5, 55-60 cmbs	139	30.2	17.5	2.1	Chip of Sunray shell, ground to sharpen. This is a small piece of a larger tool;
14.5 m N of CL at 3+275, 30-40 cm bs	163	43.0	21.4	3.1	Fragment of Sunray shell which appears to have been ground along two edges, though patination makes this uncertain;
14.6 m E of CL at 34+02.5, 78 cmbs	-	32.9	19.2	1.8	Small Sunray fragment that has been cut on one edge, and ground to a sharp edge;
Chipped					
14.1 m E of CL at 34+100, 75 cmbs	74	64.3	31.0	8.0	Broken fragment of a chipped Sunray shell tool. Figure 5-27a;
1 m W of CL at 34.148.5, found on surface	-	78.8	42.5	14.5	Sunray shell, with umbo removed, and shell edge flaked. Figure 5-27b;
16.0 m E if CL at 34+041, 10-60 cmbs	-	32.1	20.0	2.3	Sunray shell fragment, roughly rectangular though it is broken on two edges. Two edges have been chipped to make an edge. Figure 5-27c;
14.7 m E of 34+064, 42 cmbs	100	20.5	8.4	0.5	Small fragment of chipped Sunray shell that appears to have been used as blade. Tiny chips were removed bifacially along the two long edges. Figure 5-27d;
14.5 m E of CL at 34.087, 58 cmbs	65	41.3	31.8	8.9	Sunray shell tool with chipped edge;
12.3 m E of CL at 34+105, 73-80 cmbs	74	15.4	12.4	0.8	Fragment of chipped Sunray shell tool;
12.3 m E of CL at 34+105, 73-80 cmbs	74	31.6	18.2	1.6	Fragment of chipped Sunray shell tool;
3.2 m E of CL at 33+919.5	76	-	-	-	Eight small fragments of Sunray shell that may have been chipped in the process of making a tool. They are quite different from the normal break pattern seen in this shell type, but it is not possible to be certain that they are shell tool "debitage";
8 m E of CL 34+003, 110 cmbs	90	34.6	13.3	1.6	Fragment of Sunray shell tool with chipped edge;
8 m E of CL 34+003, 110 cmbs	90	22.2	12.7	0.8	Fragment of Sunray shell tool with chipped edge;
5 m E of CL at 34+212.8, 65 cmbs	133	23.8	13.6	0.6	Thin fragment of a Sunray shell with chipped semi-circles along one edge;
Lightning Whelk (<i>Busycon sinistrum</i>)					
"Cups/Scoops"					
15.6 m E of CL at 33+635, 92-110 cmbs	88	104.3	84.2	180.0	Large Lightning Whelk with the lip of the body whorl cut and smoothed, the columnella below the outer body whorl cut off and smoothed. Three impacts are evident, one on top of whorl, and two on the outer body whorl, and one of these broke a hole in the shell. This artifact is a type sometimes referred to as a "cup" or "scoop". Figure 5-28;
12.9 m E of CL at 33+905, 113-122 cmbs	53	68.4	44.0	27.4	Small Lightning Whelk shell with columnella cut short and smoothed. The edge of the body whorl may have been cut and smoothed, but it has been broken except at the end near the columnella cut;

Table 5-3. Continued...

Provenience	Feature	Length (mm)	Width (mm)	Wgt. (g)	Notes
Lightning Whelk (<i>Busycon sinistrum</i>)					
“Cups/Scoops”					
15.5 m W of CL at 34+22.5, 95-105 cmbs	141	72.0	55.7	22.0	Medium-small Lightning Whelk with edge of body whorl cut and smoothed and the end of the columnella cut off and smoothed. Most of the spire is missing;
“Gouges”					
14.5 m N of the Cl at 3+275, 30-40 cmbs	163	120.9	80.40	115.4	Large Lightning Whelk with edge of body whorl cut and smoothed and the end of the columnella shaped into a flat point. Figure 5-29a;
16.7 m E of CL at 33+656, 15 cmbs	–	64.2	38.9	12.6	Small Lightning Whelk with the edge of the body whorl cut and smoothed and the end of the columnella shaped into a point, the the adjacent edge of the body whorl thinned to a sharp edge. Though reminiscent of a gouge, the shell would have been extremely fragile, so whatever was to be gouged would have to have been very soft. Figure 5-29b;
17.5 m E of CL at 33+657, 15 cmbs	–	42.6	6.0	1.6	Lightning Whelk columnella fragment ground into a blunt point. Figure 5-29c;
11.4 m W of CL at 34+020, 65 cmbs	–	57.5	9.5	3.3	Lightning Whelk columnella fragment ground into a very sharp point. Figure 5-29d;
Backdirt near 34+015 on west side.	–	42.7	3.8	0.7	Columnella of a small Lightning Whelk, ground smooth with two sharp points. Figure 5-29e;
9 m W of CL at 34+006.5, 140 cmbs	137	81.9	8.5	4.3	Lightning Whelk columnella fragment ground into a flat point, which subsequently broke off. The proximal end was also ground to a blunt, rounded point. Figure 5-29f;
11.7 m W of CL at 34+017, 79 cmbs	148	51.3	10.1	3.5	Lightning Whelk columnella with most of body whorl removed, the distal end smoothed. At the distal end a small part of the body whorl has been left, and smoothed into a sort of “tooth” adjacent to the smoothed scoop of the distal columnella, probably to use as a type of gouge. Figure 5-29g;
10.2 m E of CL at 33+560, 143 cm bs	15	91.0	50.8	29.8	Medium-small Lightning Whelk with edge of body whorl cut and smoothed and the end of the columnella shaped into a point;
32.2 m E of CL at 34.151, 73 cmbs	99	26.2	9.4	1.5	Lightning Whelk columnella fragment ground into blunt point;
14.6 m E of CL at 34+139.2, 25-32 cmbs	114	17.3	6.0	0.3	Columnella fragment of a very small Lightning Whelk with the distal end ground to a blunt, flat point;
7.4 m E of CL at 33+546.5, 152 cmbs	125	34.1	8.8	1.8	Broken tip of a smoothed Lightning Whelk columnella;
7.4 m E of CL at 33+546.5, 152 cmbs	125	43.1	8.9	3.5	Broken tip of a smoothed Lightning Whelk columnella;
11.7 m W of CL at 34+017, 79 cmbs	148	42.0	15.5	1.5	Small Lightning Whelk, with most of body whorl removed and the columnella ground to a sharp point;
11.7 m W of CL at 34+017, 79 cmbs	148	62.6	12.7	1.5	Columnella fragment of a small Lightning Whelk with the distal end ground to a blunt, flat point;
14.5 m N of the Cl at 3+275, 30-40 cmbs	163	52.8	27.3	7.8	Small Lightning Whelk with edge of body whorl cut back but the edge along the suture not smoothed. The end of the columnella was cut at an angle and smoothed;
14.5 m N of the Cl at 3+275, 30-40 cmbs	163	44.0	20.0	3.7	Small Lightning Whelk with end of the columnella was cut at an angle and smoothed;
16.5 m E of CL at 33+653.5, found in backdirt	–	60.7	9.8	3.3	Lightning Whelk columnella fragment ground into a flattened, somewhat rounded point;
16.5 m E of CL at 33+653.5	–	43.3	7.5	1.5	Columnella of a small Lightning Whelk, with distal end smoothed into a scoop shape;
16.5 m E of CL at 33+653.5, 15 cmbs	–	45.3	7.5	2.1	Lightning Whelk columnella fragment ground into a sharp point;

Table 5-3. Continued...

Provenience	Feature	Length (mm)	Width (mm)	Wgt. (g)	Notes
Lightning Whelk (<i>Busycon sinistrum</i>)					
Scrapers and smoothers					
16.7 m S of CL at 3+233, 25 cmbs	147	148.5	53.1	62.2	Large Lightning Whelk with entire outer body whorl removed and edge near suture smoothed to an edge, possibly for use as a scraper. The columnella does not appear to have been modified. Figure 5-30a;
Backdirt near 34+004.5 on west side	–	94.3	52.9	40.3	Medium-sized Lightning Whelk with outer body whorl removed and edge near the suture smoothed, possibly for use as a scraper. Figure 5-30b;
Backdirt near 27 m E of CL at 34+192.5 on west side	–	139.0	58.3	116.4	Large Lightning Whelk with outer body whorl cut off and the edge smoothed to the suture line, which is highly polished. The end of the columnella has been sharpened as well, though it is broken. Figure 5-30c;
13 m E of CL at 33+856, 75 cmbs	91	72.7	59.7	39.5	Medium-small Lightning Whelk with outer body whorl removed and edge near the suture smoothed, possibly for use as a scraper. The columnella was shaped into a point, but has been broken;
15.5 m W of CL at 34+22.5, 95-105 cmbs	141	72.0	37.3	15.4	Small Lightning Whelk with outer body whorl removed and edge near the suture smoothed, possibly for use as a scraper;
11.7 m W of CL at 34+017, 79 cmbs	148	43.2	34.6	7.8	Small Lightning Whelk with about half the outer body whorl removed and the edge near the suture smoothed, possibly for use as a scraper. Most of this has broken away. The columnella is broken near the expansion of the body whorl and somewhat smoothed after the break;
14.5 m N of the Cl at 3+275, 30-40 cmbs	163	36.9	28.4	8.4	Small Lightning Whelk with edge of body whorl cut back about half way around and smoothed, and the edge along the suture smoothed. The end of the columnella was cut and smoothed;
19.7 m N of CL at 3+412.6, 97 cmbs	170	44.5	21.0	3.6	Small Lightning Whelk, with body whorl partially removed, and edge of shell at suture smoothed. Columnella is shaped into a flat point;
4 m W of CL at 33+485, in backfill of old trench	–	73.6	43.5	31.8	Medium-small Lightning Whelk with outer body whorl removed and edge near the suture smoothed;
“Adze Blanks”					
8 m E of CL 34+003, 90-110 cmbs	90	70.1	57.6	32.0	Possible “adze” blank, removed from Lightning Whelk body whorl, but no further work done;
15.5 m W of CL at 34+22.5, 95-105 cmbs	141	174.4	78.4	122.8	Section of the body whorl of a very large Lightning Whelk that has been removed, and one section cut off. There is no evidence that the piece itself was used as a tool;
Eastern Oyster (<i>Crassostrea virginica</i>)					
Ground					
15.0 m E of CL at 33+763, 60-67 cmbs	89	–	–	–	An oyster shell fragment ground to sharpen. Would have been fragile, so could only be used on soft materials. Figure 5-31a;
Perforated					
11.8 m W of CL at 34+034.5, 52-55 cmbs	149	136.7	65.2	69.5	A large oyster with a hole cut near the umbo. The hole is triangular, and the upper layers were cut in a triangular shape from the outside, while the lower levels were cut in a smaller triangle, measuring 11.3 mm long and 6.4 mm at the base. There may be slight evidence of wear at the point of the triangle. Figure 5-31b;
12.1 m W of Cl at 34+83.2, 46-50 cmbs	155	74.6	48.0	20.3	Perforated oyster shell, with hole 10.1 mm long, and some evidence of wear on one corner. Figure 5-31c;
14.6 m E of CL at 33+589, 110 cmbs	84	63.2	46.6	12.6	Perforated oyster shell, with a small (4.2 mm) round hole cut into it. Figure 5-31d;
14.5 m E of CL at 34+136	114	58.0	39.4	9.9	Perforated oyster shell, with hole 11.4 mm in diameter, broken across hole;
14.5 m E of CL at 34+136	114	79.2	46.8	19.5	Perforated oyster shell, with hole 14.8 mm in diameter, broken across hole;
14.5 m E of CL at 34+136	114	58.8	31.1	9.5	Perforated oyster shell, with hole 13.3 mm in diameter, broken across hole;
14.5 m E of CL at 34+136	114	53.3	38.8	19.4	Perforated oyster shell, with hole 12.3 mm in diameter, broken across hole;
14.5 m E of CL at 34+136	114	63.6	35.6	10.1	Perforated oyster shell, with hole 15.6 mm in diameter, broken across hole;

Table 5-3. Continued...

Provenience	Feature	Length (mm)	Width (mm)	Wgt. (g)	Notes
Eastern Oyster (<i>Crassostrea virginica</i>)					
Perforated					
14.5 m E of CL at 34+136	114	49.1	34.3	6.4	Perforated oyster shell, with hole 12.2 mm in diameter, broken across hole;
14.5 m E of CL at 34+136	114	41.8	35.5	5.1	Perforated oyster shell, with hole 8.9 mm in diameter, broken across hole;
Southern Quahog (<i>Mercenaria campechiensis</i>)					
11.7 m W of CL at 34+017, 79 cmbs	148	47.4	45.7	16.6	Quahog shell shaped into a rough triangle by grinding. The entire shell has considerable patination and some encrustation. The grinding and shape of this shell is reminiscent of "adzes" made from large Lightning Whelk body whorls. Figure 5-32a;
Banded Tulip (<i>Fasciolaria tilium</i>)					
12.9 m E of CL at 33+905, 113-122 cmbs	53	15.0	5.0		Banded tulip shell with a rectangular section cut from outer body whorl. Figure 5-32b;
Florida Horse Conch (<i>Pleuroploca gigantea</i>)					
13.02 m E of CL at 33+942, 130 cmbs	-	276.6	127.8	508.8	Moderate sized Horse Conch. Opened in 2 body whorls presumably to access meat. Most of outer body whorl removed by bashing (most likely for "adzes". Side opposite where the two whorls were opened has been used for grinding something;

The specimen shown in Figure 5-26c is another type of chipped shell. It is rectangular and was chipped along one edge and at least part of an adjacent edge, almost forming a right angle where they meet. The fourth tool Figure 5-26d is made on a thin, blade-like shell flake rather than the entire shell valve. It has bifacial chipping along opposing long edges. Remnants of the outer shell are visible along one end.

Modified Lightning Whelk Shells

The several modified Lightning Whelk tools can be divided into four principal groups: cups/scoopers (n=4), gouges (n=7), scrapers/smoothers (n=3), and adze blanks (n=2).

"Cups/scoopers". These artifacts were made by cutting off the columnella at the point where the last body whorl narrows to form the outer edge of the siphonal canal. It is usually supposed that these shells were used as cups, scoops, or dippers. The edge of the body whorl would be too fragile to use this tool as a digger.

Four cup/scoops were identified during the project (see Table 5-3), of which one (Figure 5-27) is an almost complete example. This artifact was deliberately broken, with three areas of impact damage, one of which broke through the outer body whorl.

"Gouges". Most modified Lightning Whelks appear to have been used as some type of gouge or awl. The

end of the siphon canal of a Lightning Whelk is semi-circular and has fairly sharp edges, which would make a reasonable gouge unmodified, however unmodified it would be very susceptible to breakage. Most gouge-like modifications on the siphon canal consist of grinding away the outer edge, leaving the main body of the columnella, and then shaping the end into one of three types: 1) a sharp, rounded point; 2) a blunt rounded point (possibly very worn examples of #1); and c) a flatten blunt point. Sometimes these shells are left almost unmodified except for the end of the columnella and the rim of the outer body whorl. More often, the most or all of the body whorls are removed, leaving only the columnella, which is smoothed by grinding. When this is the case, the columnellas are often bi-pointed.

The specimen pictured in Figure 5-28a is an example of a shell with minimal modification. The rim of the body whorl has been cut evenly and smoothed and the remaining portion of the columnella smoothed into a fairly flat, oval point. This artifact resembles the "cup" type, such as Figure 5-27, except for the shaping of the columnella, and may have been used for multiple purposes. The specimen in Figure 5-28b is another minimally modified shell with the columnella shaped into a sharp point and with the outer edge of the siphon canal ground to a sharp edge. This tool would have been a very effective gouge or awl, but only for very soft materials, as the shell is quite fragile. The specimen in Figure 5-28c is a broken tool with all the body whorls removed, leaving only a slight ridge, which was then smoothed carefully, and with the end shaped into a

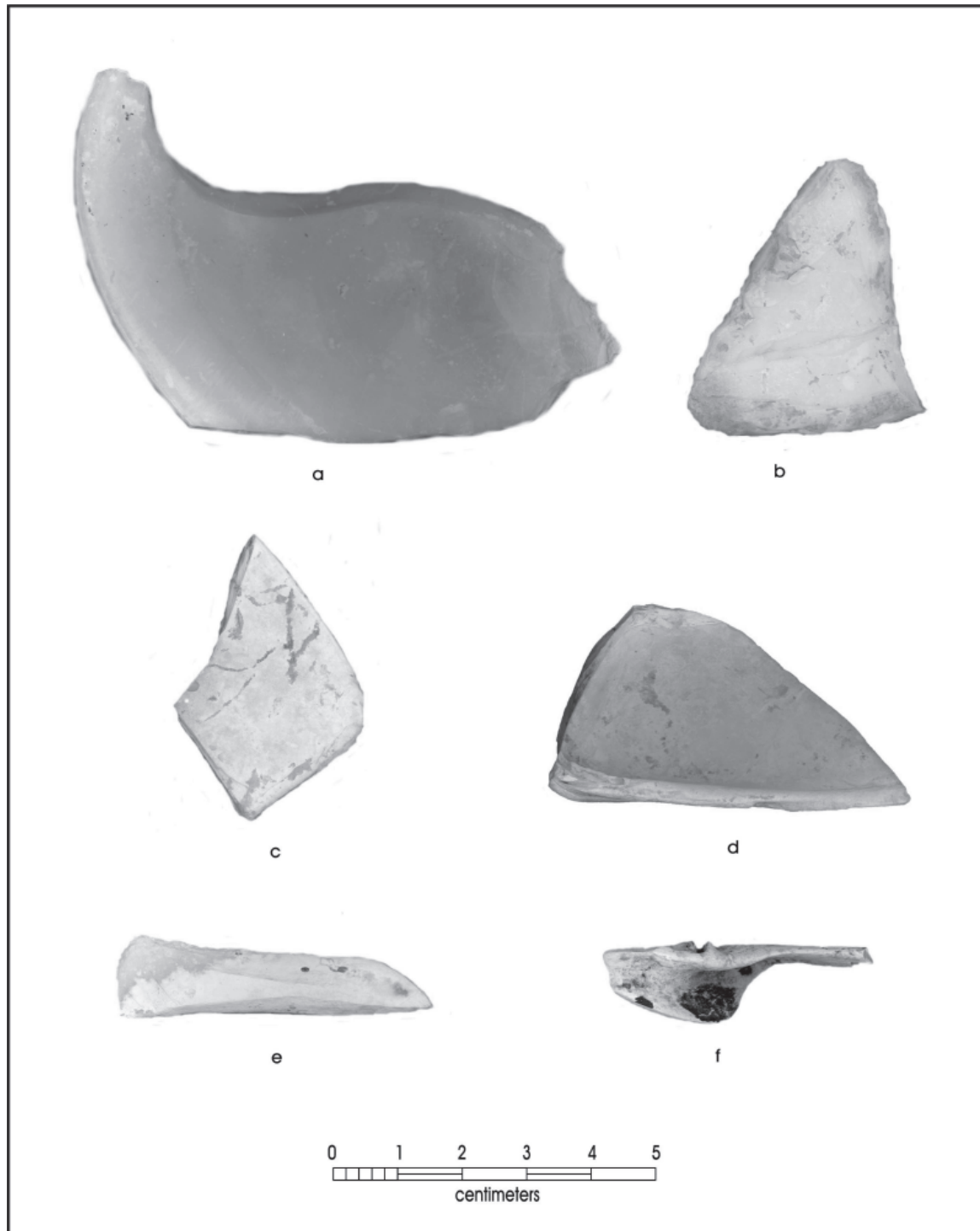


Figure 5-25. Ground shell tools: a) large tool modified along one edge; b) triangular tool, Feature 76; c) rectangular tool, Feature 65; d) tool made from umbo of *Sunray*, Feature 98; e) blade-like tool, Feature 170; f) long thin point ground from hinge plate of *Sunray*, Feature 151.

blunt point. The item in Figure 5-28d is a complete tool with body whorls removed and the columnella smoothed. The end is a sharp point. The tool in Figure 5-28e is an example of a bi-pointed columnella specimen with sharp points on both ends. The tool shown in Figure 5-28f is also

bi-pointed, but in this case some of the body whorl was left unsmoothed, with the proximal end shaped into a blunt point and the distal end shaped into a flat point that has broken. The specimen shown in Figure 5-28g is a fragment of a columnella that has been shaped into a blunt point,

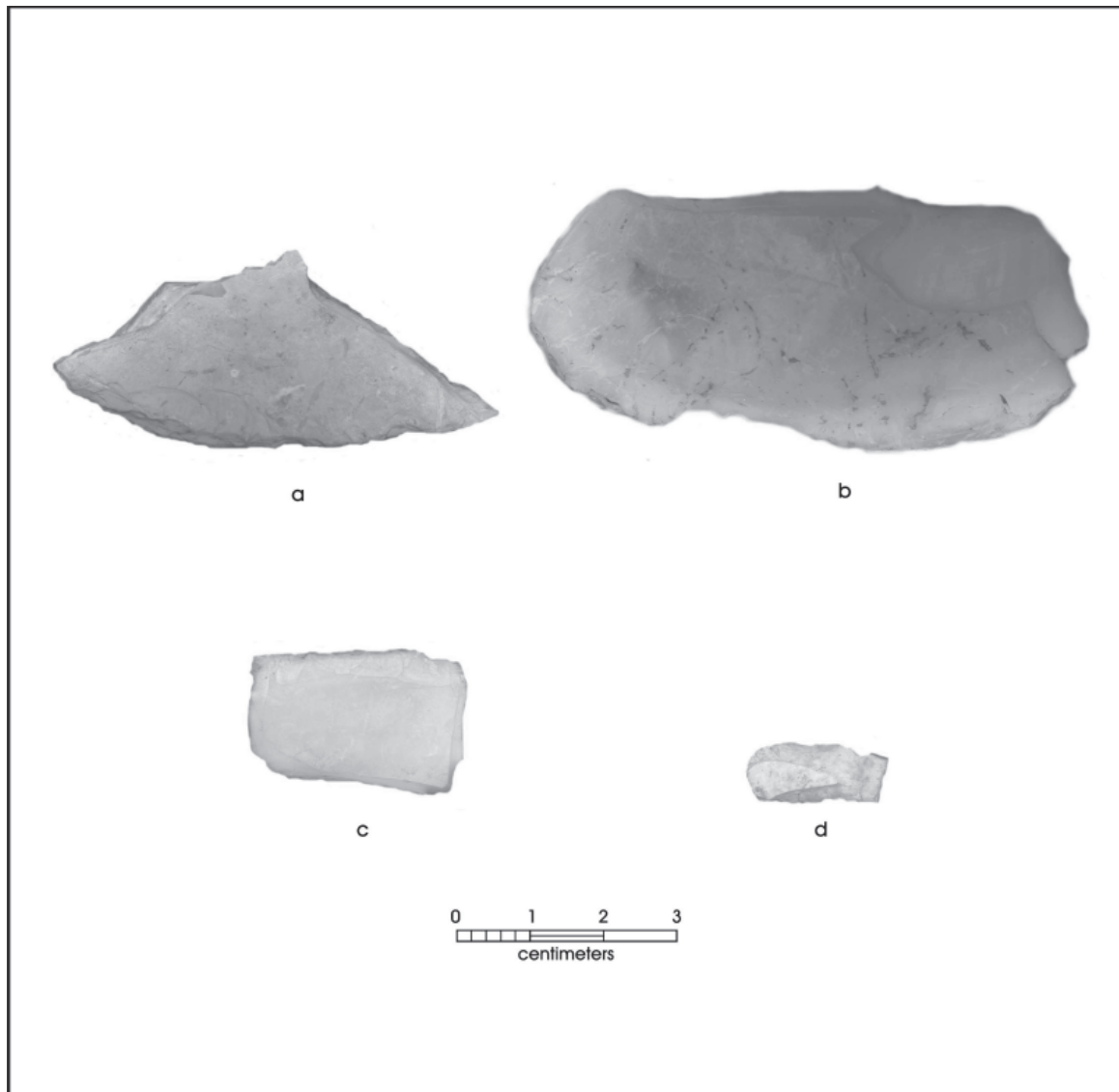


Figure 5-26. Chipped Sunray shell tools: a) fragment of tool, Feature 74; b) fragment of tool from surface of North Dune; c) modified shell fragment; d) small bifacially flaked blade-like tool, Feature 100.

with a remnant of the outer edge of the siphon canal left in place and shaped to form a sort of “tooth” near the point.

Scrapers and “Smoother”. Lightning Whelk scrapers are made by removing a section of outer body whorl, leaving a “ledge” near the suture, which is then smoothed into a relatively sharp edge. These tools would have been held by the columnella and pulled toward the worker. “Smoother” is a name given to shells with the outer body whorl removed to the suture and exhibiting heavy polish along the suture. In some cases, it appears that a tool starts out as a scraper but is later used as a smoother as the edge is broken away. Many shells modified to act as scrapers or smoothers have also had the end of the columnella modified to a point. These tools come in many sizes.

The tools shown in Figures 5-29a and 5-29b are scrapers. The columnella of Figure 5-29a does not appear to have been altered. Note that the right side of the specimen in Figure 5-29b shows a broken body whorl while the left side shows the shaping associated with a scraper. The specimen in Figure 5-29c is classified as “smoother”, with no fragment of the body whorl remaining and the junction at the suture showing considerable smoothing and polish. The end of the columnella has been shaped into a point.

“Adze Blanks”. Artifacts referred to as “adzes” are not uncommon in sites in the Coastal Bend area. These are sections of shell, usually removed from the outer body whorl of a large Lightning Whelk or Horse Conch (and very occasionally other species, see below), and shaped into a

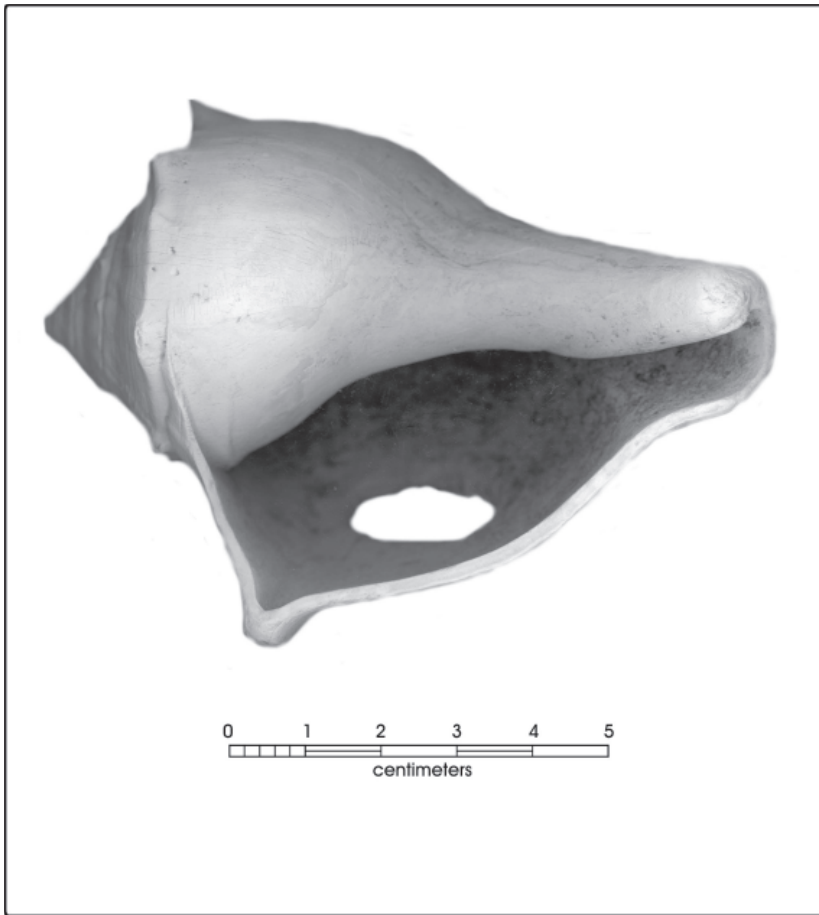


Figure 5-27. A Lightning Whelk shell modified to form a cup or scoop, Feature 88.

roughly triangular shape with the wide end ground to make an edge. While the exact purpose of these tools is unknown, we will continue to call them “adzes” as this is the term used in the literature (Ricklis 2004:Figure 5.14). An adze “blank” is a section of large gastropod that shows evidence of having been deliberately removed from the body whorl, but does not show any other modifications. Two such artifacts were recovered during the project (Table 5-3), they are not shown due to their fragmentary condition.

Modified Oyster Shells

One fragment of oyster shell was recovered with clear grinding along two edges (Figure 5-30a). The purpose of this artifact is unknown, as oyster shell has a tendency to fragment and is not often used for tools. It may be part of an object

shaped by grinding for another purpose, such as decoration.

Ten perforated oysters were recovered from the site. These are common artifacts along the coast (Ricklis 2004:Figure 5.6). It is generally assumed that the oysters serve as weights, possibly for nets. The specimens shown in Figures 5-30b through 5-31d are examples of these artifacts.

Modified Quahog Shell

A single modified quahog shell was identified (Figure 31a). It was shaped into a rough triangle by grinding, and appears to be an example of an “adze”, though most adzes recorded in the area are made from the body whorl of Lightning Whelks. Rather than use the nature thinning at the ventral edge of the shell, the adjacent edge was ground to serve as the bit for the “adze”. The entire shell has considerable patination and some encrustation. The shell had been in water after the death of the animal, as it exhibits a large number of worm holes and two large holes probably made by sea fauna.

Modified Banded Tulip Shell

A single Banded Tulip shell was modified by the cutting out of a rectangular piece of the outer body whorl (Figure 5-31b). The removed section of shell is only 5 by 15 mm and the purpose it may have been put to is unknown, but it may have been used for decoration.

Modified Florida Horse Conch Shell

One medium-sized Florida Horse Conch was recovered that shows use for a number of purposes (Figure 5-32). Two body whorls in the spire had been broken open, presumably to help remove the conch meat. The entire outer body whorl was removed, apparently in two sections, probably for use in manufacturing “adzes”. In addition, the protuberances along the top of the last body whorl have been worn smooth, evidence that the shell was used for grinding.

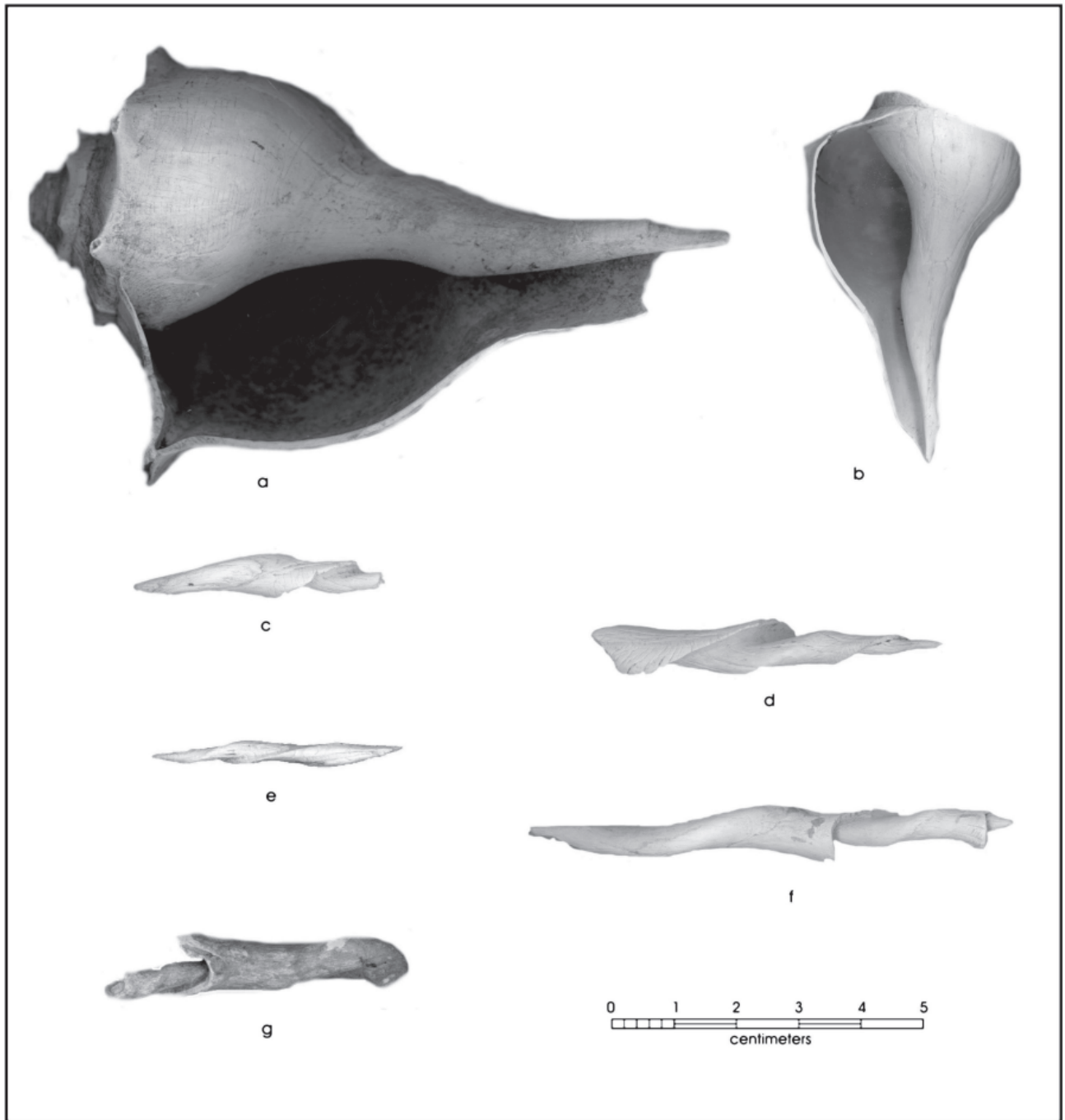


Figure 5-28. Lightning Whelk Gouges: a) collumella shaped into flat point, Feature 163; b) small whelk with collumella shaped into a point; c) collumella fragment ground into a blunt point; d) collumella tool shaped into a sharp point; e) bi-pointed collumella tool; f) collumella fragment with a broken flat point on one end and a blunt point on the other; g) collumella fragment with shaped "tooth".

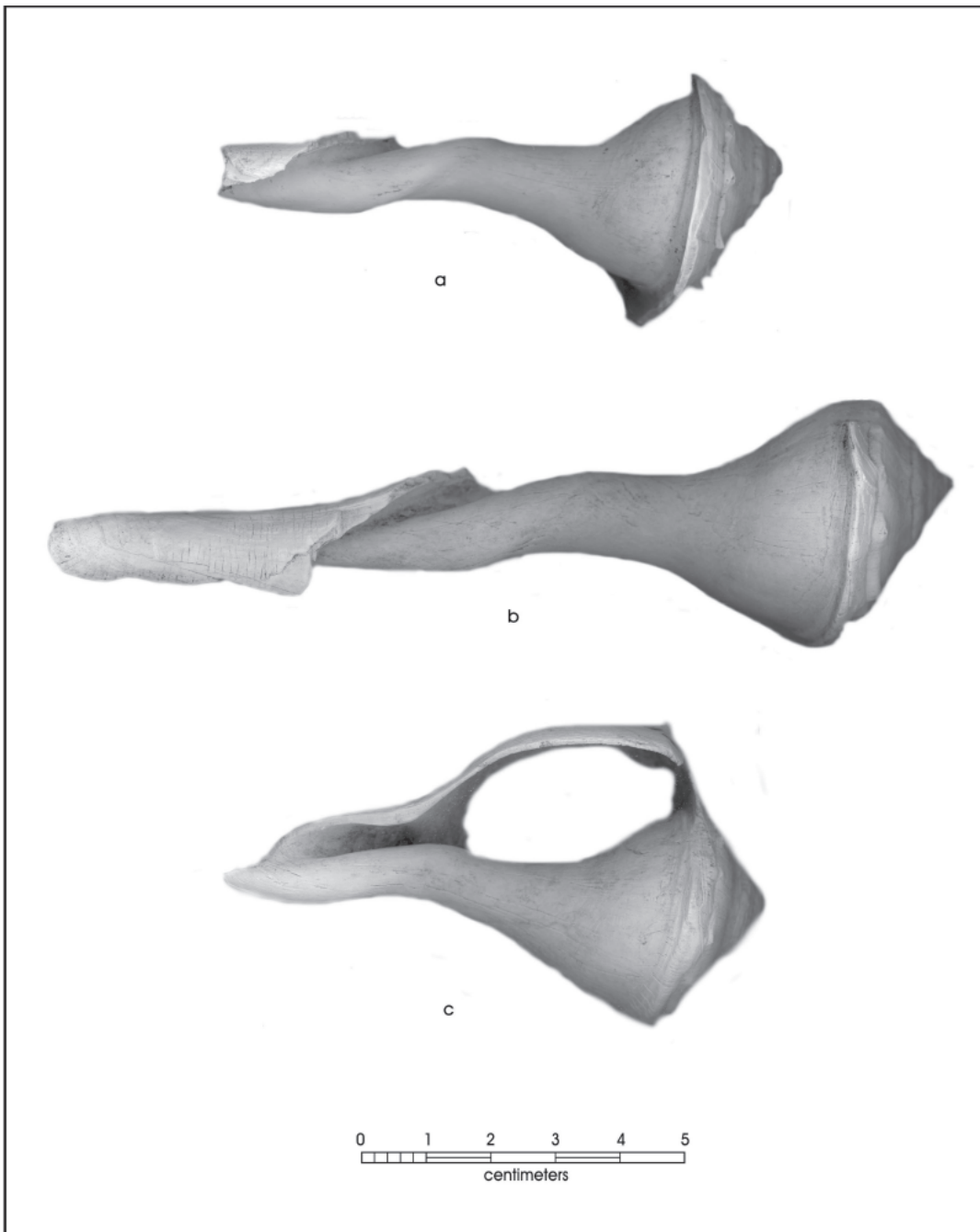


Figure 5-29. *Lightning Whelk* scrapers and “smoothers”: a-b) scrapers; c) a “smoother”.

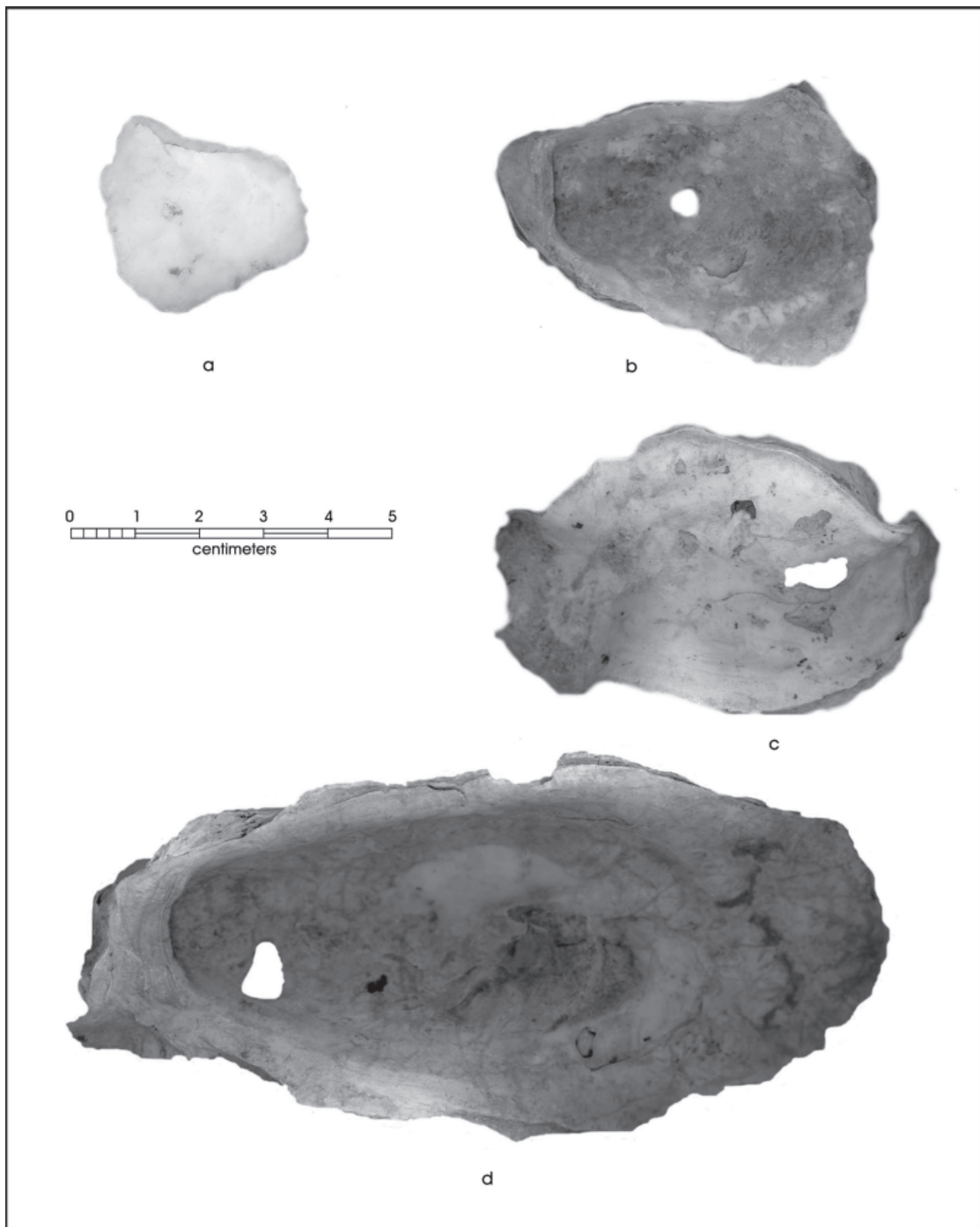


Figure 5-30. Modified Easter Oysters: a) a fragment of oyster shell with two ground edges; b-d) perforated oysters.

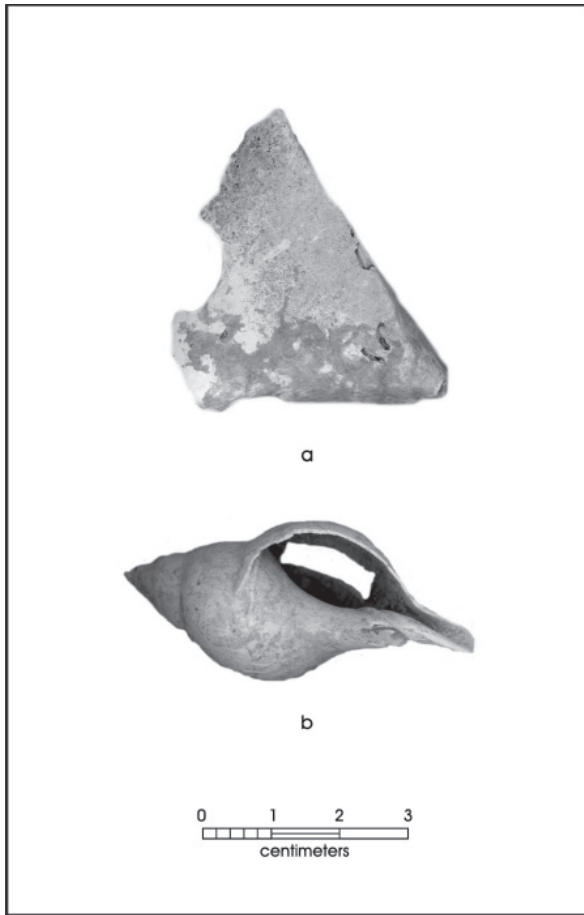


Figure 5-31. *Other modified shells: a) an “adze” made from a quahog shell; b) a Banded Tulip shell with a rectangular section removed.*

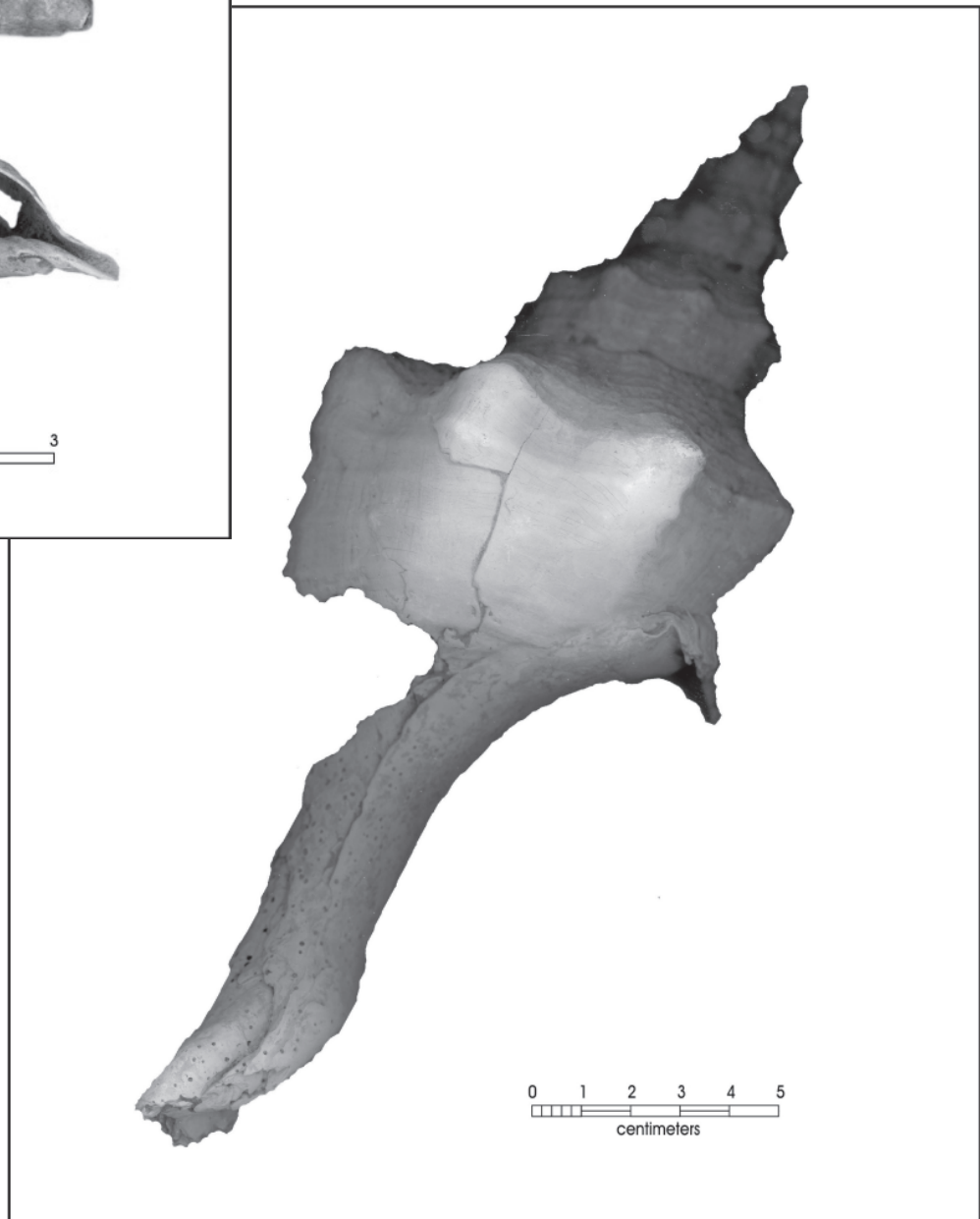


Figure 5-32. *A Florida Horse Conch showing extensive modifications.*

Chapter 6: Radiocarbon Dates, Feature Descriptions and Occupation Zones

Barbara M. Meissner and Raymond P. Mauldin

This chapter provides a summary of the radiocarbon results, along with descriptive data on 164 features, and 11 occupation zones, documented at 41NU2 during the monitoring project. The focus on the radiocarbon discussion is on comparing dates derived from 40 samples, 20 of which were derived from marine shell and 20 of which represent assays from charcoal or organic sediment. The paired samples were primarily derived from individual features or areas of dense accumulations of cultural debris termed occupation zones. Our comparison shows that in 11 of the 20 pairings, the corrected, calibrated date ranges derived from shell do not overlap at 2 sigma with date ranges provided by charcoal or organics from the same provenience. Nevertheless, a comparison of the mid-points of the two sigma ranges for the individual dates shows, with one exception, a strong relationship. We conclude that marine shell can provide accurate, precise dates in the absence of other dating materials. Appendix G supports the radiocarbon section. Following the radiocarbon results, we summarize the 164 features documented on the project. We present a feature typology based primarily on feature content (e.g. shell clusters, charcoal and burnt clay clusters) and overall form. Finally, the chapter concludes with a discussion of the 11 occupation zones identified at the site during the monitoring.

Radiocarbon Dates

Fifty-eight radiocarbon dates were acquired during this project. This number includes 12 humate dates associated with coring activities, 40 dates from features, and 6 dates from other contexts. This section concentrates on the 46 dates derived primarily from features and occupation zones. The 12 humate dates will be discussed in the context of the paleoenvironmental data (Appendix F). Forty of the 46 dates considered here represent paired samples of marine shell and organics selected from the same provenience, occupation zone or feature. The paired samples were designed to assess the reliability of marine shell as a dating material, as well as to provide basic site chronology.

The original sample selection strategy had three principal goals. First, we wanted to date as many features as possible with multiple samples to determine whether features represented relatively short-term use or repeated reuse. Second, date the tops and bottoms of thick organically enriched and poor zones to define ages and accumulation rates. Finally, whenever possible, we wanted to date paired charcoal and shell samples

to examine the correspondence between the dates. This strategy resulted in a preliminary list of 110 samples. Given that the costs associated with such a large sample size were prohibitive, TxDOT-ENV and CAR developed an alternate sampling strategy with less ambitious goals that preserved some of the original goals, but at a reduced cost. The final dating strategy consisted of the selection of samples from two sequences of features. One sequence was to be selected from the north dune (n=3) and one from the south dune (n=3). In addition, paired samples were selected from 20 features or locations, one sample consisting of charcoal (n= 20) and the other of marine shell (n = 20). This revised strategy allowed for investigation of the correspondence between charcoal and shell dates, as well as a coarse-grained dating of the stratigraphic sequences noted in each of the two dunes.

Results

Beta Analytic of Miami, Florida ran all samples. Appendix G provides details on all 58 samples. Table 6-1 summarizes the results of the 46 samples considered here. The CAR sample number in Table 6-1 (e.g., **371**) is an abbreviated version of the sample number referenced in Appendix G (e.g., 41NU2-**371**). All samples have been corrected and calibrated, and in the following discussion, we will primarily use the two sigma range, or the midpoints of that range.

Using the midpoint of the two sigma span, the radiocarbon dates in Table 6-1 range from 5930 B.P. (5990-5870 B.P.) on shell from Feature 49 to around 905 B.P. (1010-800 B.P.) on shell from Feature 54. The early Feature 49 shell date (Table 6-1; Sample # 426) is roughly 2400 years older than any other date from the site. In addition, the date is about 4150 years older than the paired sample, run on organics, from this feature (Table 6-1; Sample # 381), which has a date range of 1870 to 1700 B.P. It is likely that the shell date is unreliable. We have chosen to eliminate sample 426 from additional consideration.

Figure 6-1 shows the distribution of the remaining 45 radiocarbon dates plotted at a two sigma range. Examination of the figure suggests that the site was minimally occupied from around 3500 B.P. to about 900 B.P. Note that there are two periods without dates during this sequence. There is a small gap just after 1500 B.P. and a roughly 400 year gap between 3400 B.P. and just after 3000 B.P. To the degree that these dates on 23 features and three occupation zones reflect overall site use, it appears that the occupation of 41NU2 began after 3500

Table 6-1. Radiocarbon Results from Features and Occupation Zones

CAR Sample #	Feature	Provenience	Depth (cm bgs)	Dune	Material Used	2 Sigma Cal. (B.P.)	Conv. RC Age (B.P.)
371	2	11.5 m W of CL at 33+509	30	S	Charcoal	1170-960	1140 +/-40
436	2	11.5 m W of CL at 33+509	30	S	<i>Macrocallista nimbosa</i>	980-840	1370 +/-40
418	7	11.7 m of CL at 33+475	245	S	Charcoal	3470-3350	3190 +/-40
416	10	11.0 m of CL at 33+590	90	S	Charcoal	2710-2350	2440 +/-40
437	16	10.2 m of CL at 33+549	96	S	<i>Busycon sinistrum</i>	1160-960	1510 +/-40
375	16	10.2 m of CL at 33+549	96	S	Organic sediment	1300 - 1180	1320 +/-40
430	18	25.5 m N of CL at 3+325	75	N	<i>Macrocallista nimbosa</i>	1690 - 1510	2030 +/-40
376	18	25.5 m N of CL at 3+325	75	N	Organic sediment	1940 - 1740	1910 +/-40
420	31	12.00 m E of CL at 33+894	86	out	<i>Crassostrea virginica</i>	1770 - 1560	2090 +/-40
377	31	12.00 m E of CL at 33+894	86	out	Organic sediment	1820 - 1600	1780 +/-40
380	46	11.17 m E of CL at 34+022	70	N	Charcoal	2050 - 1880	2010 +/-40
428	46	11.17 m E of CL at 34+022	70	N	<i>Busycon sinistrum</i>	2340 - 2150	2590 +/-40
381	49	11.33 m E of CL at 34+021.5	59	N	Organic sediment	1870 - 1700	1840 +/-40
426	49	11.33 m E of CL at 34+021.5	59	N	<i>Macrocallista nimbosa</i>	5990 - 5870	5540 +/-40
419	54	12 m E of CL at 33+813.7	56	out	<i>Busycon sinistrum</i>	1010 - 800	1390 +/-40
382	54	12 m E of CL at 33+813.7	56	out	Organic sediment	1290-1160	1290 +/-40
425	63	11.0 m E of CL at 34+059.5	57	N	<i>Macrocallista nimbosa</i>	2720-2480	2850 +/-40
384	63	11.0 m E of CL at 34+059.5	57	N	Organic sediment	3070-2800	2830 +/-50
417	81	14.95 m E of CL at 33+564	80	S	Charcoal	1350 - 1270	1390 +/-40
438	84	14.5 m E of CL at 33+584.5	110	S	<i>Crassostrea virginica</i>	2290 - 2060	2500 +/-40
387	84	14.5 m E of CL at 33+584.5	110	S	Charcoal	2740 - 2460	2510 +/-40
431	103	14.7 m E of CL at 34+019	80	N	<i>Macrocallista nimbosa</i>	1970 - 1800	2270 +/-40
391	103	14.7 m E of CL at 34+019	80	N	Organic sediment	2300-1990	2120 +/-40
393	108	14.6 m E of CL at 34+091.8	37	N	Organic sediment	1410 - 1300	1470 +/-40
440	110	14 m E of CL at 33+572	155	S	<i>Crassostrea virginica</i>	3630 - 3440	3640 +/-40
396	110	14 m E of CL at 33+572	155	S	Charcoal	3630 - 3440	3300 +/-40
412	113	14.4 m E of CL at 34+111.7	98	N	Charcoal	2130-1930	2060 +/-40
397	115	14.8 m E of CL at 34+138	70	N	Organic sediment	2770 - 2720	2610 +/-40
429	115	14.8 m E of CL at 34+138	70	N	<i>Macrocallista nimbosa</i>	2950 - 2760	3090 +/-40
434	33/121	14.1 m E of CL at 33+901	120	out	<i>Crassostrea virginica</i>	2540-2300	2700 +/-50
398	33/121	14.1 m E of CL at 33+901	120	out	Charcoal	2730-2360	2470 +/-40
433	122	14.7 m E of CL at 33+907	110	out	<i>Busycon sinistrum</i>	2010-1840	2310 +/-40
399	122	14.7 m E of CL at 33+907	110	out	Charcoal	2320-2060	2180 +/-40
432	123	14.5 m E of CL at 33+915	85	N	<i>Busycon sinistrum</i>	1050-900	1430 +/-40
400	123	14.5 m E of CL at 33+915	85	N	Charcoal	1170-960	1140 +/-40
415	137	8 m W of CL at 34+152	140	N	Charcoal	2730-2350	2450 +/-50
422	153	12.1 m W of CL at 34+067.3	44	N	<i>Busycon sinistrum</i>	1790 - 1580	2100 +/-40
403	153	12.1 m W of CL at 34+067.3	44	N	Organic sediment	2300-2000	2140 +/-40
423	155	12.1 m W of CL at 34+083.2	48	N	<i>Crassostrea virginica</i>	2310 - 2110	2540 +/-40
405	155	12.1 m W of CL at 34+083.2	48	N	Organic sediment	2480 - 2340	2370 +/-40
427	n/a	14.8 m E of CL at 34+079	60	N	<i>Macrocallista nimbosa</i>	1880-1590	2110 +/-40
424	n/a	14.8 m E of CL at 34+079	50	N	<i>Macrocallista nimbosa</i>	1880-1700	2200 +/-40
388	n/a	14.8 m E of CL at 34+079	50	N	Organic sediment	1930 - 1730	1900 +/-40
389	n/a	14.8 m E of CL at 34+079	60	N	Charcoal	1980 - 1820	1940 +/-40
439	n/a	14.0 m E of CL at 33+567	124	S	<i>Crassostrea virginica</i>	2280-2030	2480 +/-40
373	n/a	14.0 m E of CL at 33+567	124	S	Organic sediment	2360-2200	2310 +/-40

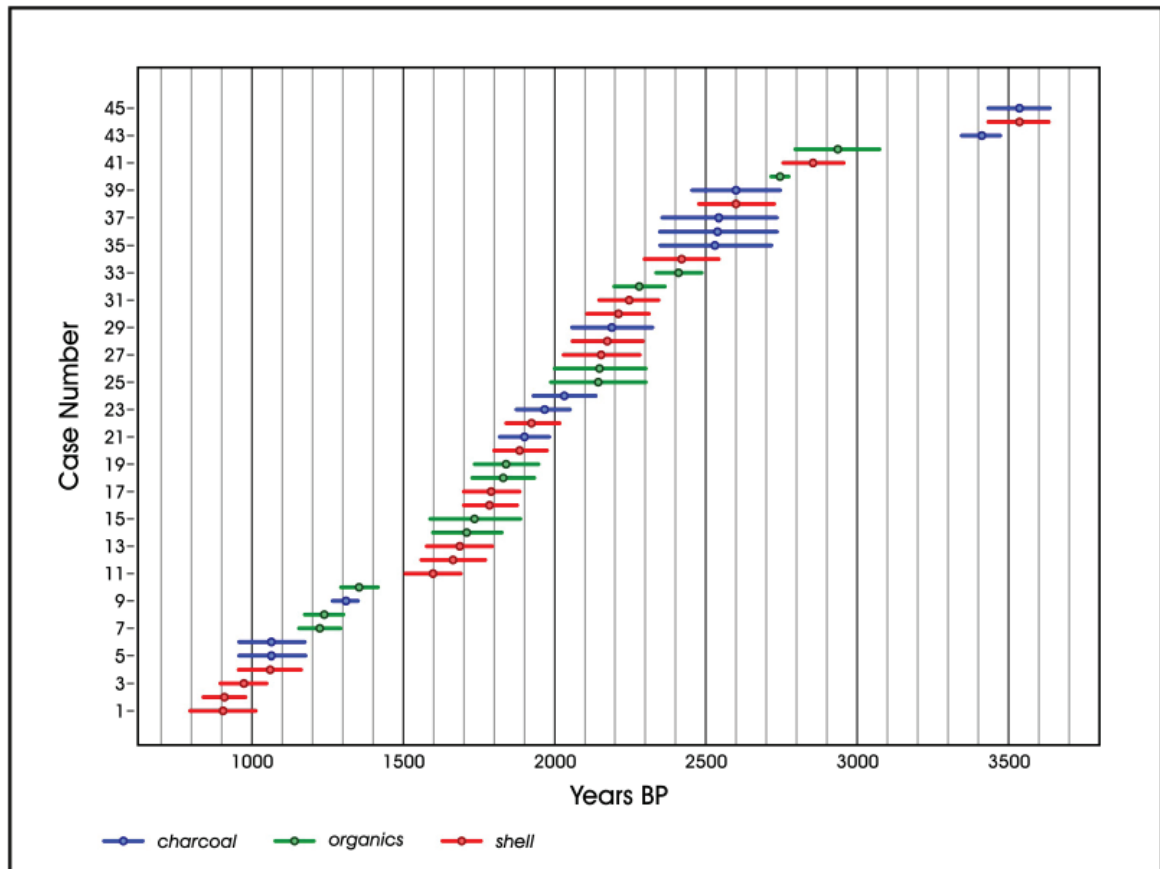


Figure 6-1. Radiocarbon dates associated with monitoring at 41NU2. Plotted at 2-sigma.

B.P., as represented by a cluster of dates from Features 7 and 110. Most occupation of the site falls between 3000 and 1600 B.P., with no evidence of site use after 900 B.P.

Figure 6-2 shows the distribution of dates from the North (top) and South dunes (bottom). Both dunes show roughly similar occupation patterns. The primary difference between the dunes appears to be that all three pre-3000 B.P. dates (Table 6-1; Samples 418, 440, and 396) are from the South dune.

Assessing Marine Shell Dates

One of the primary interests of the revised dating strategy noted above was to investigate the utility of radiocarbon dates on marine shell for age determination. Researchers have debated the accuracy of shell dates for decades (e.g., Anderson and Libby 1951; Berger et al. 1964; Gillespie and Polach 1979; Head 1982; Johnson 1955; Keith et al. 1964). Early comparisons of charcoal and shell dates from the same context frequently produced discrepancies of several hundred years, with uncorrected shell dates often being significantly older than charcoal dates. Marine shellfish produce shell from both organic carbon, derived primarily from diet, as well as

inorganic carbon derived from ocean water. Radiocarbon analysis has focuses primarily on the inorganic component, essentially monitoring the ^{14}C activity of the ocean water in which the shell formed. One of the major sources of error in the accuracy of marine shell dates is, then, related to delays in exchanges between atmospheric CO_2 and ocean carbonates, as well as impacts related to upwelling of older ocean waters and those related to fractionation (see Berger et al. 1966; Mangerud 1972; Stuckenrath 1977; R.E. Taylor 1987). Shell carbonates are also particularly susceptible to post-depositional recrystallisation, a process that can dramatically alter the dating accuracy (see Goslar and Pazdur 1985).

While the accuracy of marine shell dates will remain suspect, researchers have made significant advances understanding the sources of errors and developed corrections for those impacts (see R.E. Taylor 1987). Differences in fractionation are well understood and can be corrected, and shell dates are routinely calibrated against a marine calibration curve. Figure 6-3 presents a section of that curve, along with the associated atmospheric curve (Ramsey 2003). A comparison of the two curves shows that for the same radiocarbon determination (Y-Axis), the calibrated date

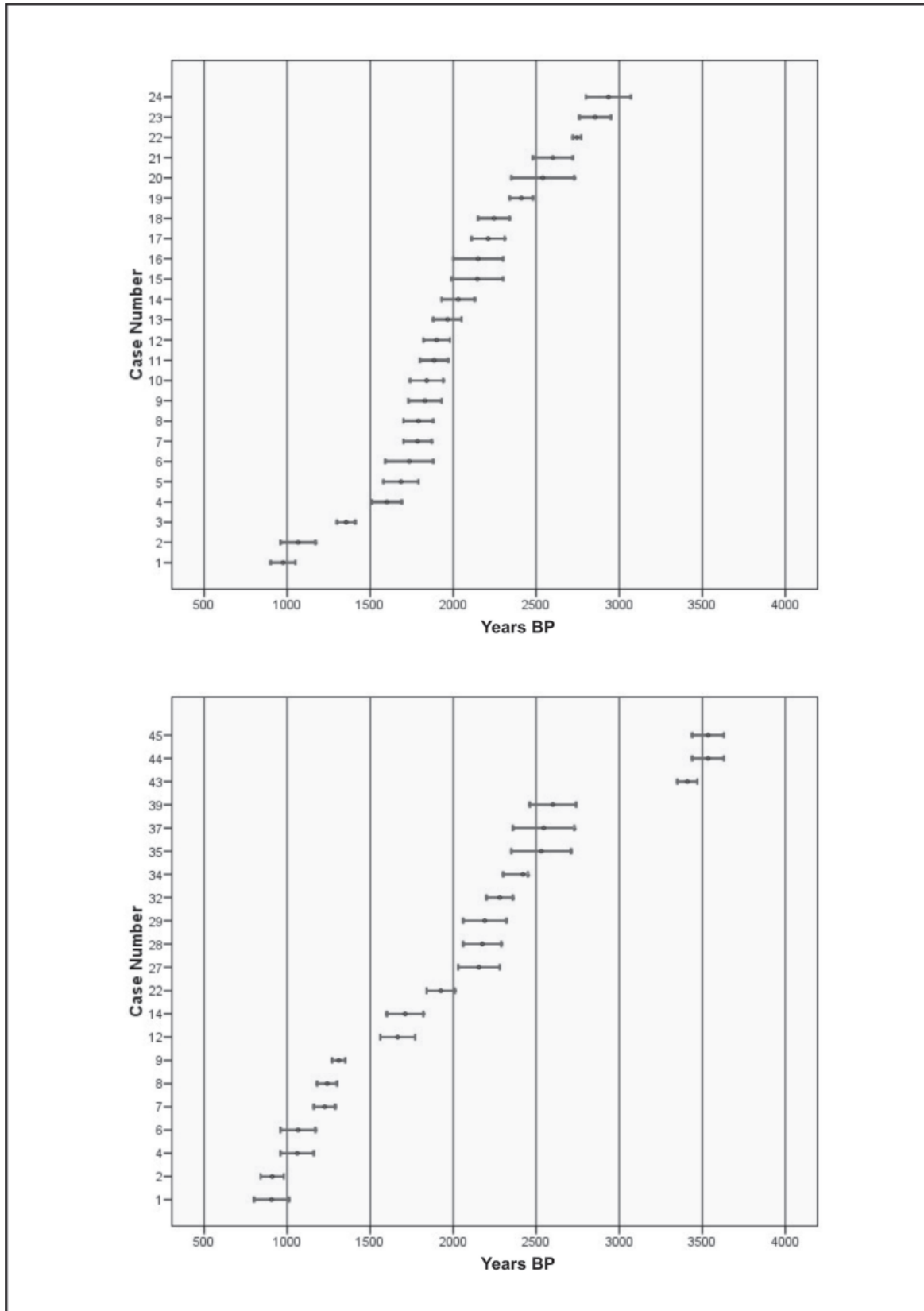


Figure 6-2. Top - Radiocarbon dates from the North Dune. Bottom - Radiocarbon dates from the South Dune and Beaumont Ridge. Plotted at 2-sigma.

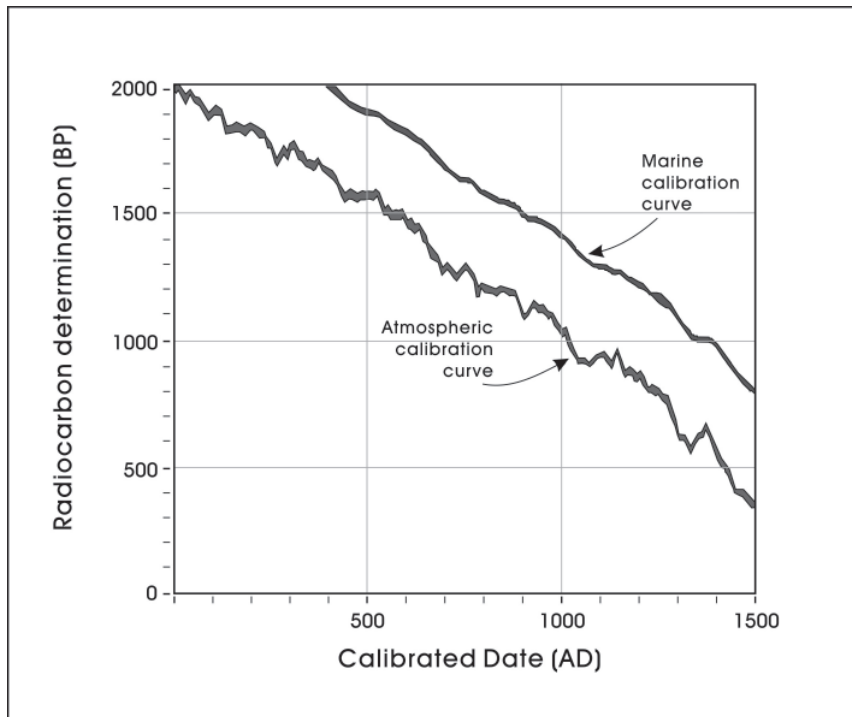


Figure 6-3. Comparison of marine (blue) and atmospheric (red) calibration curves from A.D. 1 to A.D. 1500.

(X-Axis) differs by about 400 years. In addition, the Figure 6-3 curves also show at least one advantage to dating shell. Because of the effects of the oceanic reservoir, the marine calibration curve lacks the wiggles that are common in the curve used for terrestrial samples. Shell dates should produce calibrated dates that are more precise than dates derived from charcoal of other organics that rely on the atmospheric calibration.

For the 41NU2 comparisons, 20 charcoal or organic sediment samples were paired with 20 marine shell samples collected from the same feature or occupation zone. Table 6-2 presents the 2-sigma ranges for dates on organics and the associated shell counterpart for a given provenience. Assuming that the charcoal or organic dates are an accurate reflection of the true date, and using the 2-sigma range, only 45% (9 of 20) of the shell date ranges overlap their charcoal or organic counterparts (Table 6-2). Frequently, the corrected, calibrated shell dates are more recent than the date obtained from their charcoal/organic sediment counterpart.

Table 6-2. Dates for Matched Samples (Organics - Shell) at 2-Sigma, 41NU2

Feature or FS #	Organics or Charcoal 2 Sigma Cal. (B.P.)	Marine Shell 2 Sigma Cal. (B.P.)	Overlap Between Date Ranges	Marine Shell Species
2	1170-960	980-840	YES	<i>Macrocallista nimbosa</i>
16	1300 - 1180	1160-960	NO	<i>Busycon sinistrum</i>
18	1940 - 1740	1690 - 1510	NO	<i>Macrocallista nimbosa</i>
31	1820 - 1600	1770 - 1560	YES	<i>Crassostrea virginica</i>
46	2050 - 1880	2340 - 2150	NO	<i>Busycon sinistrum</i>
49	1870 - 1700	5990 - 5870	NO	<i>Macrocallista nimbosa</i>
54	1290-1160	1010 - 800	NO	<i>Busycon sinistrum</i>
63	3070-2800	2720-2480	NO	<i>Macrocallista nimbosa</i>
84	2740 - 2460	2290 - 2060	NO	<i>Crassostrea virginica</i>
103	2300-1990	1970 - 1800	NO	<i>Macrocallista nimbosa</i>
110	3630 - 3440	3630 - 3440	YES	<i>Crassostrea virginica</i>
115	2770 - 2720	2950 - 2760	YES	<i>Macrocallista nimbosa</i>
121/33	2730-2360	2540-2300	YES	<i>Crassostrea virginica</i>
122	2320-2060	2010-1840	NO	<i>Busycon sinistrum</i>
123	1170-960	1050-900	YES	<i>Busycon sinistrum</i>
153	2300-2000	1790 - 1580	NO	<i>Busycon sinistrum</i>
155	2480 - 2340	2310 - 2110	NO	<i>Crassostrea virginica</i>
FS 71	2360-2200	2280-2030	YES	<i>Crassostrea virginica</i>
FS 270	1980 - 1820	1880-1590	YES	<i>Macrocallista nimbosa</i>
FS 271	1930 - 1730	1880-1700	YES	<i>Macrocallista nimbosa</i>

The Table 6-2 results suggest that the 2-sigma date range on shell does not provide an accurate match with the charcoal/organic sediment date at 41NU2. However, note that some shell species, such as *Crassostrea virginica*, have relatively good correspondence, with a success rate of four overlaps in six comparisons (66.7%). Conversely, only in 1 of 6 (16.7%) cases did the date ranges of *Busycon sinistrum* overlap with the range for the paired organic sample. While sample sizes are small, and while additional comparisons will be necessary, these data suggest that relatively good success rates might be anticipated by focusing on shell from *Crassostrea virginica*.

Finally, Figure 6-4 considers the relationship between paired samples in a slightly different way. Here we use the midpoint of the 2-sigma spread and plot the corrected, calibrated shell dates (Y-Axis) against the midpoints for the organic sediment or

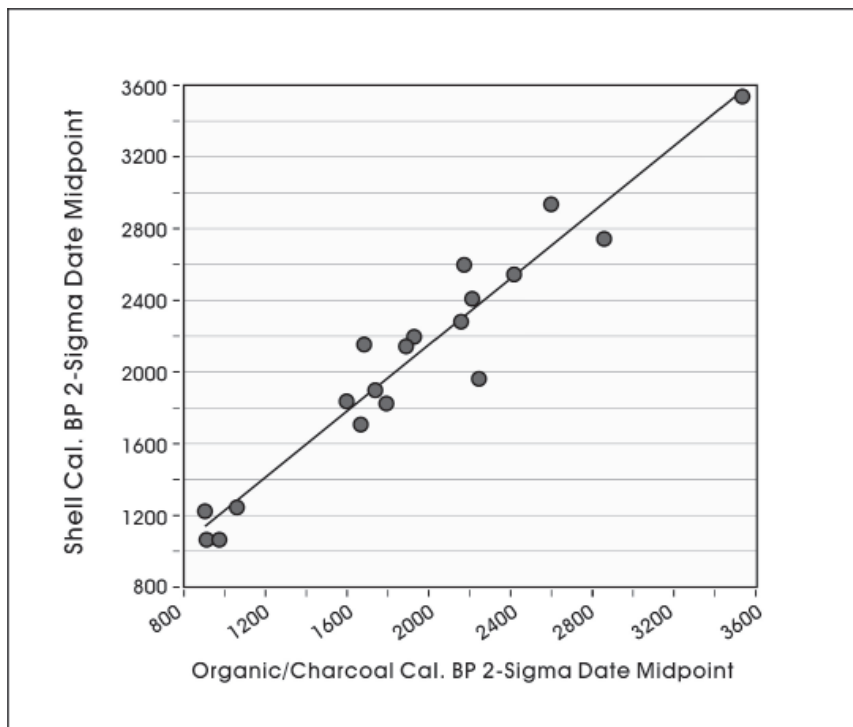


Figure 6-4. Plot of mid-point of the 2-sigma range for paired marine shell and organic samples from 41NU2.

charcoal dates (X-Axis). We eliminated one matching pair (Feature 49, samples 381 and 426, Table 6-1) which formed a significant outlier. Clearly, the figure shows a strong, linear relationship for the remaining 19 pairs. Pearson's correlation coefficient for the 19 cases is .965. Shell midpoints are, on average, about 160 years more recent than corresponding charcoal dates.

The systematic difference shown in Figure 6-4 may hint at local corrections that need to be made in the overall marine curve. However, the charcoal/organic sediment dates may also be overestimating the true feature age. We lack any details on the particular type of charcoal or organic sediment dates, and it is certainly possible that long-lived tree species make up a significant component of the samples dated. While old wood is unlikely to account for differences on the order of 160 years in the coastal region, as our paired samples lack any information on species it is probable that the true dates of feature use may be more recent than the charcoal/organic dates.

Summary

Radiocarbon dates from features and occupation zones collected during the 41NU2 monitoring project suggest that the earliest documented occupation of the site was about 3500 B.P. Occupation seems to increase after around

3000 B.P., a time frame for the possible coalescence of the barrier islands. After about 900 B.P. we lack any radiocarbon dates, suggesting that the site was not extensively used. While comparison of shell and organic dates from the same context shows that in most cases there was no overlap between date ranges, our results do suggest that certain species of shell may be more reliable as a dating tool. In addition, there is a strong linear relationship between the midpoints of date ranges derived from paired shells and organics. If we assume that the organic dates reflect the true age of a feature, then corrections on the order of 160 years may be required for shell dates to accurately reflect feature ages. However, given that the organic samples probably contain wood of varying ages, it is probable that the organic ages overestimate the true feature age. Shell dates, then, especially from *Crassostrea virginica*, have a high probability of providing accurate dates with good precision.

Features

One hundred and seventy features were defined in the field during this project. Of these, two (Features 4 and 6) were removed from the list because of changes in the definition of feature (see Chapter 4; see also Meissner 2000:5-6). Another (Feature 30), was removed because it was clearly historic or modern. In addition, two cultural manifestations originally defined as features were later shown to be part of an occupation zone and not distinct features (Features 84 and 129). Finally, Features 33 and 121 were combined into a single feature. The remaining 164 features are described in this section. For each of the four areas of the site (South Dune, Beaumont Ridge, North Dune, Pond and Northwest Ridge), a table of basic data is presented. In these tables, the features are not considered by the sequential number given to them.

During analysis, a list of feature types was defined based on 1) content or use, such as shell clusters, hearths and pits; 2) in some cases, a size category; 3) other characteristics that were found in multiple cases. For instance, many hearths did not seem to be associated with large amounts of shells, while others of the same type were. Another example is a type of shallow pit with a flat bottom, completely filled with shells, most of which were small

and consisted of a single species. A list of Feature Types with a definition for each is provided in Table 6-3. A brief description of each Feature Type listed in Table 6-3 is given below, followed by a listing of features found in each zone. Features that are simply scatters of shell, small surficial hearths, or other common and non-distinctive types will be listed in tabular form. Only if the feature presents unique characteristics or has been carbon dated, will a more detailed description be completed within the zone segments. Feature locations are shown in Figures 6-5 though 6-9.

Feature Types Found at 41NU2

Table 6-3 presents a list of Feature Types with definitions. Here a more detailed description of the most common and least distinctive types of features will be presented, with examples.

Shell Concentrations

Shell concentrations were clusters of marine/estuarine and/or land invertebrate shells, of varying density, usually with well-defined borders, and without evidence of a hearth or a pit. During Phase I, only concentrations containing charcoal were considered features in order to limit the number of features that had to be excavated. Shell concentrations were the most common feature, constituting 42.1% (69/164) of the features recorded at 41NU2. These concentrations were frequently mixtures of species, but occasionally consisted almost exclusively of a single species. Some shell concentrations were not homogenous, showing clusters of a particular shell surrounded by a more generalized scatter of different species.

Shell concentrations varied widely in size, from about 25 cm to as much as 5 or 6 m in diameter (see Table 6-3). An

Table 6-3. Feature Types Defined at 41NU2

Feature Type	Count	Definition
Shell Concentrations	69	Discrete cluster of marine or land invertebrates . May or may not include animal bone, charcoal, or burned clay flecks or nodules;
Type 1	36	Feature area is less than 10,000 cm ² ;
Type 2	19	Feature area is between 10,000 and 25,000 cm ² ;
Type 3	14	Feature area is more than 25,000 cm ² ;
Hearths	57	Discrete area in which a combination of charcoal and burned clay suggest that a fire was built at that location. May or may not have been placed in a basin dug for the purpose;
Surficial—Type 1	11	Hearth with no sign of a basin. Few or no shell associated;
Surficial—Type 2	11	Hearth with no sign of a basin. A significant number of shells or bones are closely associated;
Basin—Type 1	17	Basin is clearly visible. Few or no shells or bone associated. Less than 15 cm deep;
Basin—Type 2	11	Basin is clearly visible. A significant number of shells or bone is closely associated. Less than 15 cm deep;
Basin—Type 3	2	Basin is large and >15 cm deep. Few or no shells or bone associated;
Basin—Type 4	1	Basin is large and >15 cm deep. A significant number of shells or bone are closely associated;
Multiple	4	A group of two or more closely associated, surficial hearths, surrounded by a thin scatter of charcoal and shell fragments;
Pits	20	Areas where a pit can be defined. Most are densely packed with shells, and charcoal;
Type 1	10	Pit is no more than 15 cm deep filled with charcoal and/or shell and animal bone;
Type 2	4	Deeper than 15 cm, filled with charcoal and/or shell and animal bone;
Type 3	4	Small, shallow (3-5 cm) pit densely packed with shell most a single small shell species, most of which are whole, with a few other species, most of which are highly fragmented;
Type 4	1	Pit of any size or depth that is empty or only has a few shell or charcoal flecks in fill;
Multiple	1	A group of two or more closely associated pits surrounded by shell and charcoal;
Other Feature Types	18	Miscellaneous features other than shell concentrations, hearths and pits;
Bone fragment scatter	7	Often fairly dense but thin layers of highly fragmented or small bone with charcoal and usually a few fragmented shells. Usually consists largely of small fish bone;
Bone bed	2	Group of large, largely intact bones, with or without marine shells;
Charcoal cluster	4	Cluster of very dense charcoal with no associated burned clay, and few or no associated shells;
Rabdotus cluster	2	Dense cluster with few or no marine shells;
Burial pit	3	Pit dug to interred human remains;

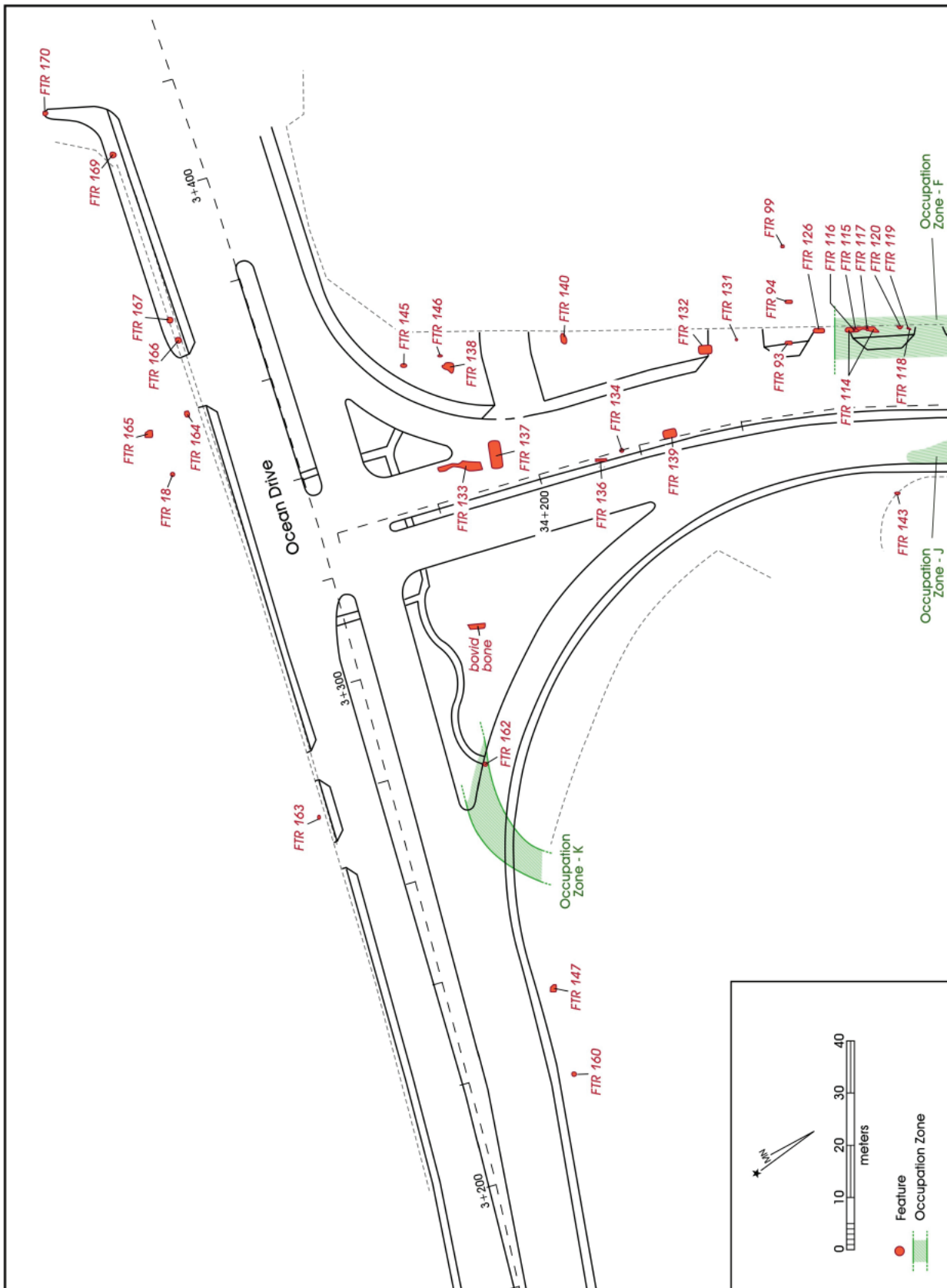


Figure 6-5. Detail of project map between 34+120 and 34+325 showing locations of features and occupation zones.

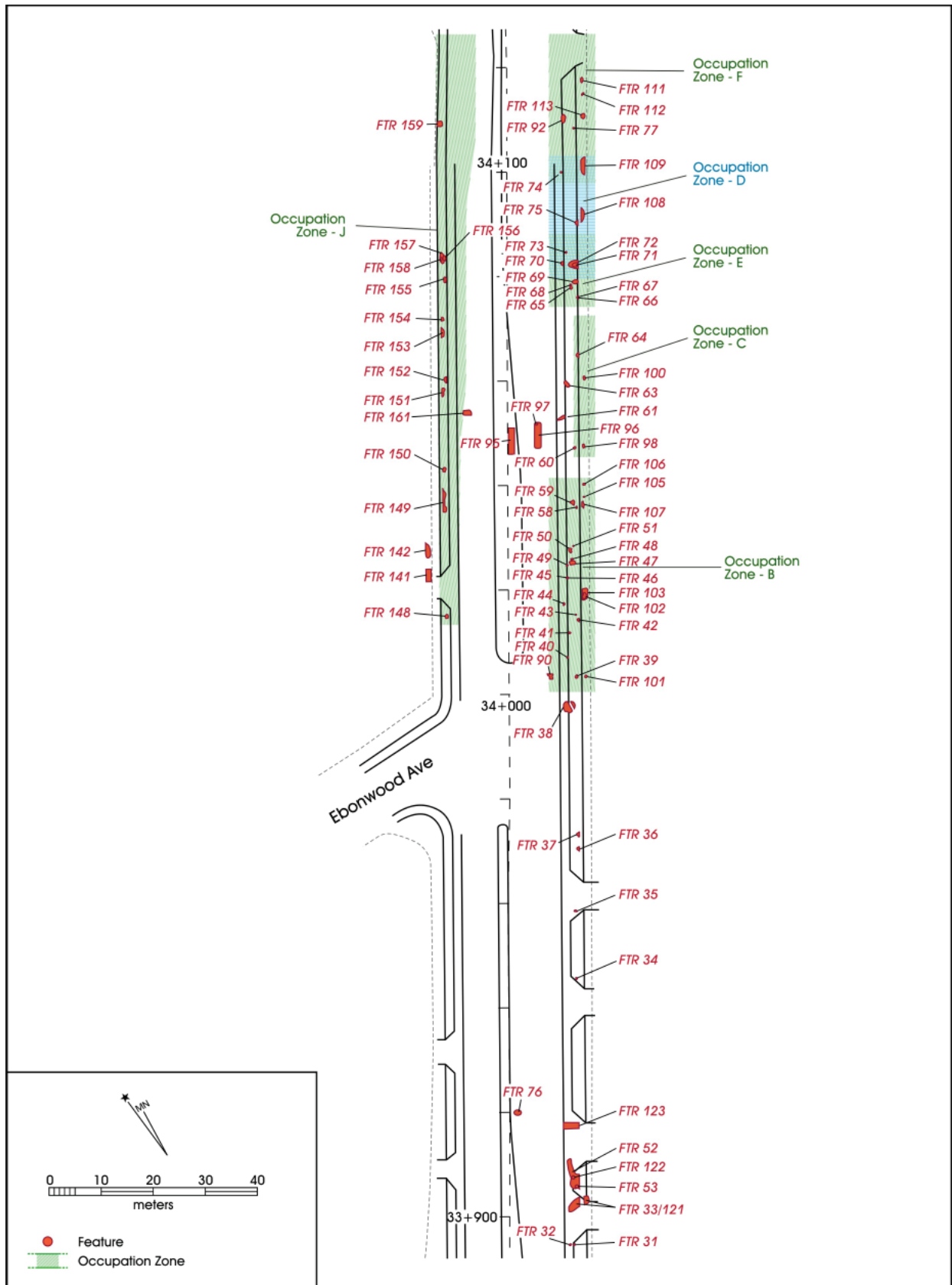


Figure 6-6. Detail of project map between 33+900 and 34+120 showing locations of features and occupation zones.

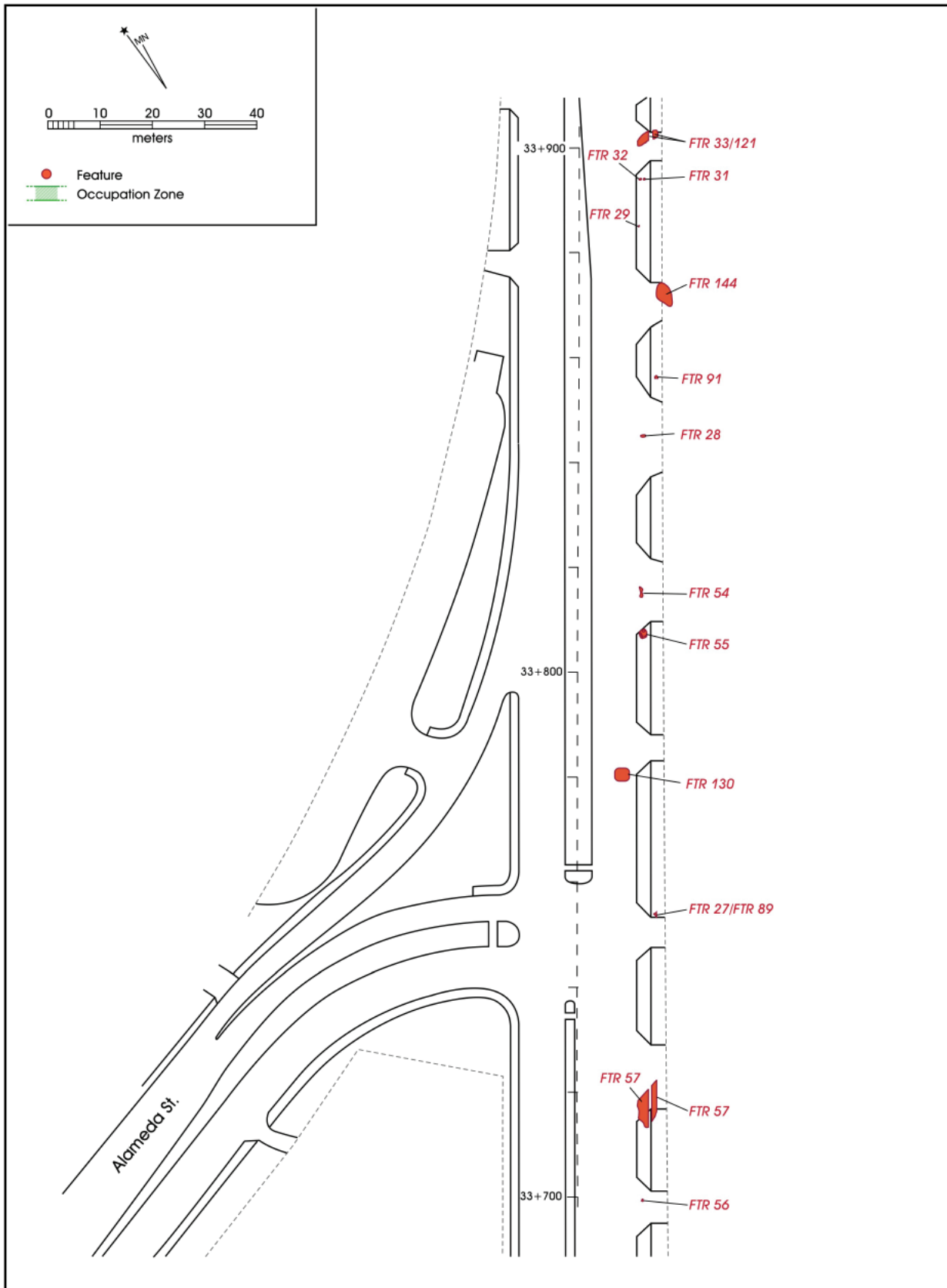


Figure 6-7. Detail of project map between 34+120 and 33+700 showing locations of features and occupation zones.

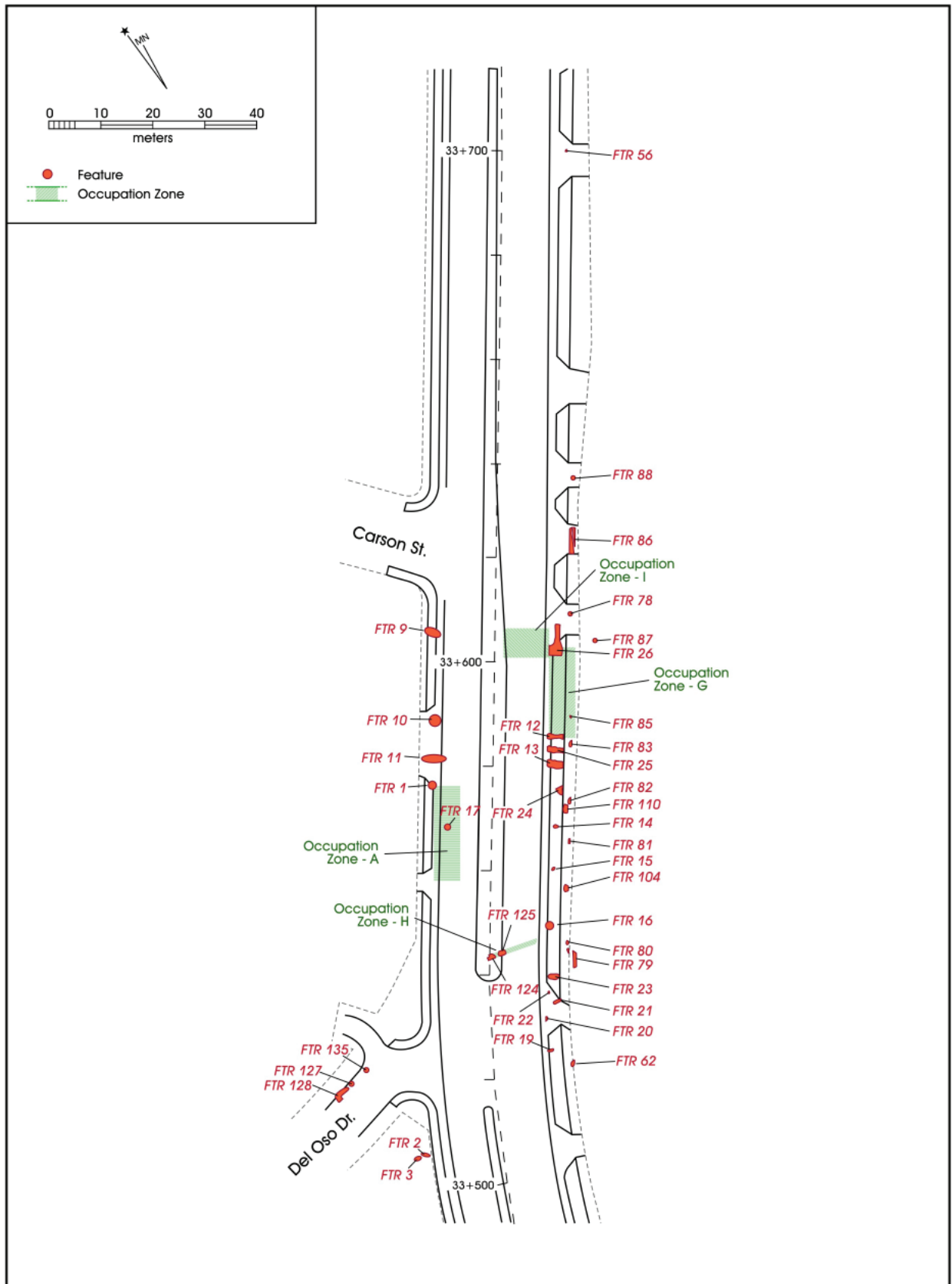


Figure 6-8. Detail of project map between 33+500 and 33+700 showing locations of features and occupation zones.

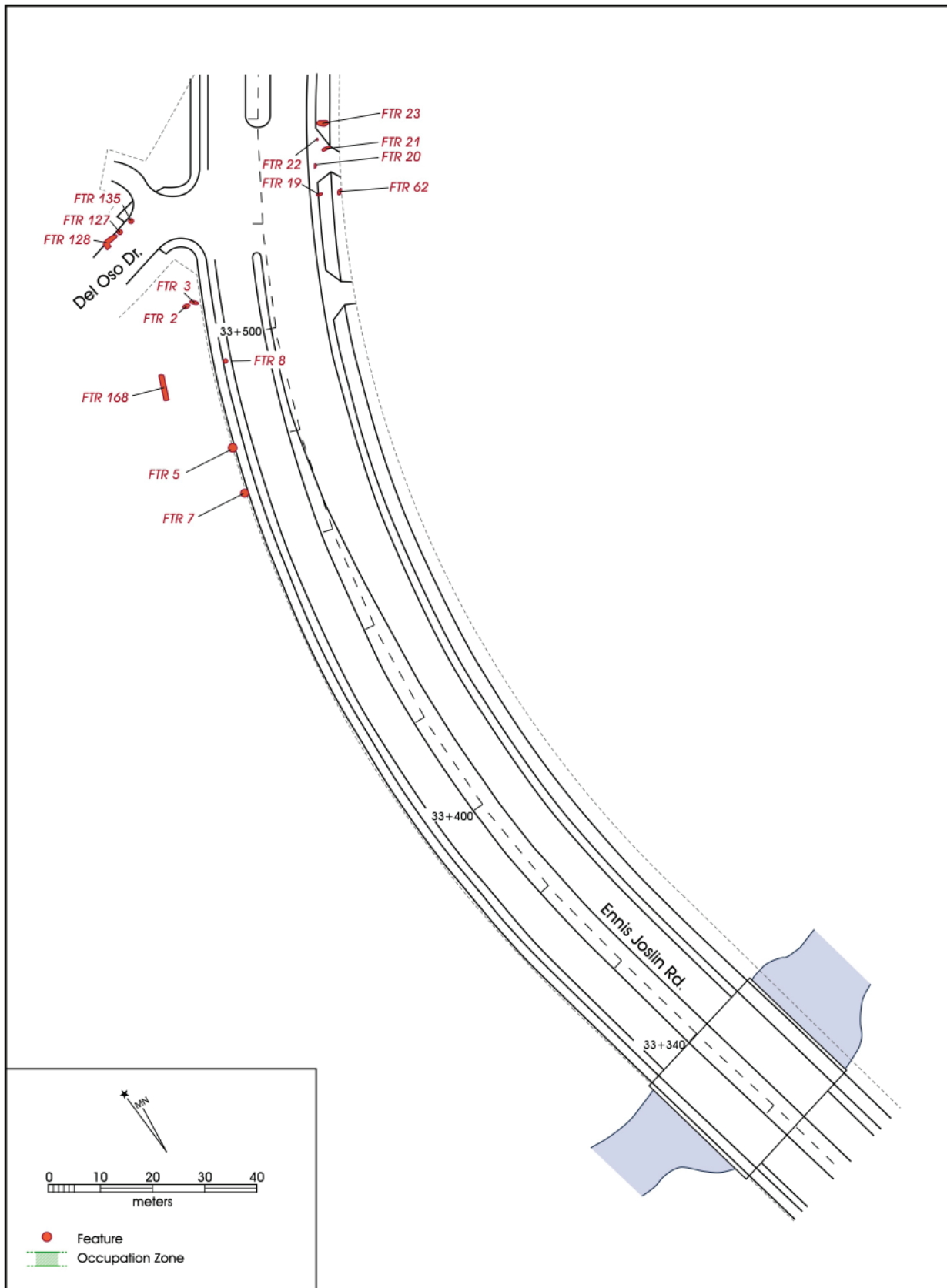


Figure 6-9. Detail of project map between 33+320 and 33+500 showing locations of features and occupation zones.

estimated area for each shell concentration was calculated using the formula: $A=\pi(X/2)(Y/2)$, where A is the area, X is the longest radius and Y is the shortest radius. Shell concentration areas varied from less than 500 cm² to more than 141,000 cm². A histogram of the areas of the shell features fell into one of three distinct groupings. There were 36 Type 1 shell concentrations, each less than 10,000 cm² constituting of 52.2 percent of the shell concentrations. Nineteen shell concentrations were between 10,000 cm² and 25,000 cm², 27.5 percent of the shell concentration features. The 14 shell concentration features measured more than 25,000 cm in area and made up 20.3 percent of the shell concentration features.

Some concentrations were relatively thin, discrete, scatters. Others were very dense (Figure 6-10). Although most were a single shell deep, a few contained several layers of shell. Most were round to oval in outline.

Hearths

Hearths were defined when it appeared that a fire was built at a location. No hearthstones were seen due to lack of rock in the environment. Hearths comprised 34.8% (57/164) of the features identified during the project (Table 6-3). A dense deposit of shells surrounded some hearths. Others had little or no associated shell. In four cases, there were multiple hearths, that is two or more distinct hearth areas in close association, usually less than a meter apart.

There were 22 hearths that appeared to have been built on the ground surface, with no pit. They comprised 41.5% of features with single hearths (Table 6-3). Hearths in a basin made up the remaining 58.5% (31/53) of features with single hearths. Surficial hearths consisted of a concentration of charcoal and burned clay that was usually only 1-3 cm deep (Figure 6-11). The surficial hearths were always small, usually no more than 30 cm in diameter, with dense charcoal and burned clay.

Basin hearths are locations where a pit was dug in the ground surface before the fire was built. Most of these pits were shallow, with 90.3 percent (28/31) of basin hearths less than 15 cm deep (Table 6-3). They were occasionally dug much deeper. Figure 6-12 shows a plan and profile of Feature 67, an example of a shallow basin hearth.

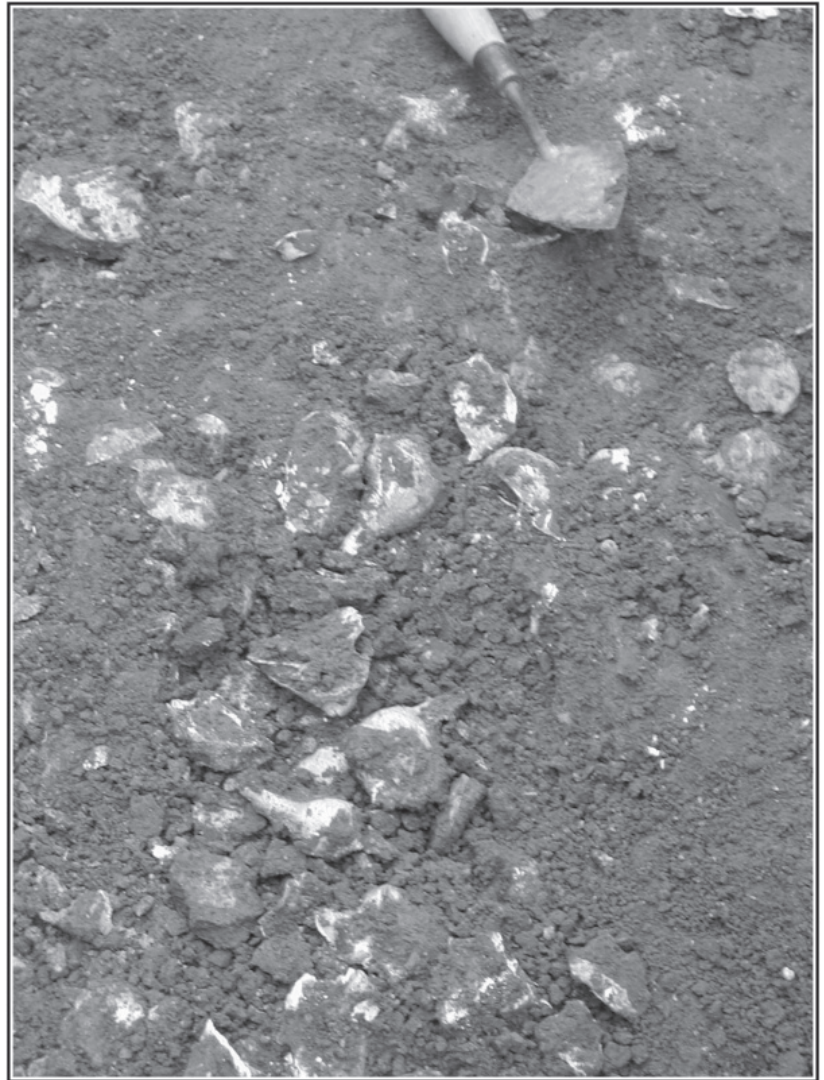


Figure 6-10. Feature 38, showing an example of a dense shell cluster.

Pits

Features were defined as pits when a basin was clearly present and it did not appear to have been dug for a hearth (e.g. Figure 6-13). There were 20 pits identified during the project, they constituted 12.2 % of the features. In many cases, pits were originally defined as small shell concentrations, and then, as the Gradall or backhoe continued digging through the feature, it could be seen that it was a pit. Feature forms were modified when this occurred.

Most pits were 5 to 15 cm deep and packed densely with shell (Figure 6-13b). Some were dug from paleosols into a light clay layer. Some pits had large quantities of animal bones among the shells. One type, Pit Type 3 (Table 6-3), consisted of small, shallow pits packed with charcoal, fragments of marine shell, and large numbers of a single species of small

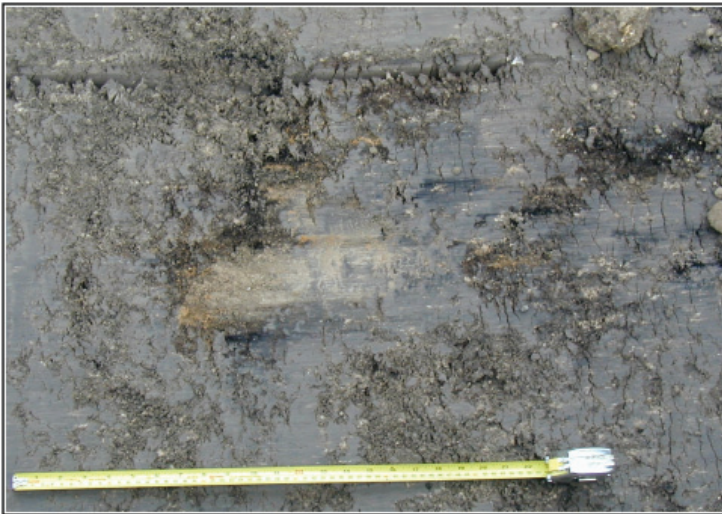


Figure 6-11. Feature 27, an example of a small surficial hearth.

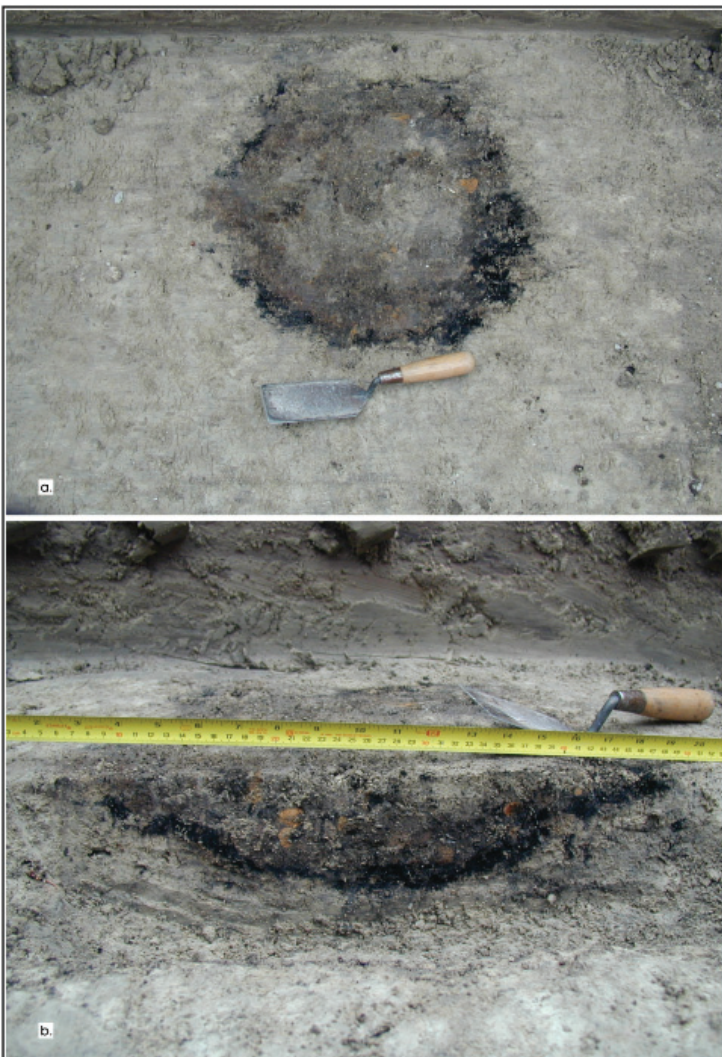


Figure 6-12. Feature 67, an example of a shallow basin hearth: a) plan view; b) profile.

shell, either Crossbar Venus (*Chione cancellata*) clams or *Rabdotus* (sp.) land snails.

Other Feature Types

Over 89% of the features fell into one of the feature types listed above. Other, less common feature types, are described below.

Bone Fragment Scatters. These features were distinctive clusters of bone, charcoal, and usually a few scattered and fragmented marine shells. They ranged in size from a tiny cluster measuring only 15 x 15 cm to a 550 x 200 cm area with a thin layer of dense bone.

Bone Bed. Bone beds are areas with many largely intact bones. There were two such features defined during the project, Features 13 and 170. Both bone bed features are described in detail below.

Charcoal cluster. These four features consisted of very dense clusters of charcoal but no evidence of burnt soil. Typically, a thin scatter of shell was associated with the charcoal.

Rabdotus cluster. This feature type does not include pits in which *Rabdotus* shells were packed. In these features hundreds, and in the case of Feature 107 (Figure 6-14), perhaps thousands of *Rabdotus* were present.

Burial Pit. A burial pit is defined as a pit in which human remains were interred. Three such pits were recorded during this project, Feature 61 (Burial 2), Feature 136 (Burial 4), and Feature 139 (Burials 6 and 7). These were discussed in the previous chapter.

Features in the South Dune

The 40 features described in this section were found on both sides of the road within the South Dune area. These 40 features represent 24.2% of the features encountered during the project. One was located in a pit dug for a manhole box, and the rest were found in trenches. Table 6-4 presents basic data on these features organized by location, south to north, and includes an in-field estimate of the kinds of shells found in each. Artifacts found in these features were discussed in the previous chapter. A analysis of the faunal remains is included in Chapter 11. Feature locations can be found on Figures 6-8 and 6-9. Samples from five features, from a cluster of shell and



Figure 6-13. Feature 102, a small pit densely packed with Lightning Whelks: a) surface of pit (note pit was dug into light sediments); b) Pit fill from Feature 102, showing dense packing of shells.



Figure 6-14. Close up of Feature 107, showing masses of Rabdotus shells.

Table 6-4. Features in the South Dune

Feature	Feature Type	Description				Estimated % of Common Shells*									
		Length (cm)†	Width (cm)†	Depth top (cmbs)‡	Bottom (cmbs)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria liliium</i>	Sunray Venus Clam <i>Macrocallyista nimbosa</i>	Southern Quahog <i>Mercenaria campechiensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.
7	Rabdotus cluster	60	30	242	245	+	+	—	—	—	—	—	—	—	+
5	Bone fragment scatter	15	15	240	245	—	—	—	—	—	—	—	—	—	+
168	Shell Concentration— Type 3	470	100	60	67	—	10	—	—	—	90	—	—	—	—
8	Shell Concentration— Type 1	40	30	45	55	—	27	—	—	—	73	—	—	—	—
2	Shell Concentration— Type 1	120	40	30	40	—	—	—	—	—	100	—	—	—	—
3	Shell Concentration— Type 1	100	80	40	45	—	—	—	—	—	100	—	—	—	—

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

* Estimates based on observation in the field. Actual counts from sampling are included in Fauna section of Chapter 11.

“+” sign means presence mentioned in notes; “—” sign means absent/not observed.

Table 6-4. Continued...

Feature	Feature Type	Description				Estimated % of Common Shells*									
		Length (cm)†	Width (cm)†	Depth top (cmbs)‡	Bottom (cmbs)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria litium</i>	Sunray Venus Clam <i>Macrallista nimbosa</i>	Southern Quahog <i>Mercenaria campechiensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.
135	Shell Concentration— Type 1	30		80	87	—	+	—	+	—	—	+	—	—	—
19	Bone fragment scatter	150	75	165	172		—	—	—	—	—	—	—	—	—
62	Hearth, Basin—Type 1	15	?	244	247		—	—	—	—	—	—	—	—	—
128	Shell Concentration— Type 1	100	100	60	74	—	—	—	—	—	—	95	—	—	—
127	Shell Concentration— Type 2	300	100	72	87	+	+	—	+	—	—	—	—	—	—
20	Shell Concentration— Type 1	75	75	160	168	10	60	—	5	—	20	+	—	+	—
21	Shell Concentration— Type 1	150	75	175	185	—	48	—	+	—	48	—	—	+	—
22	Shell Concentration— Type 2	175	178	30	40	80	8	—	—	—	—	8	+	—	—
23	Shell Concentration— Type 2	200	100	175	178	20	57	+	10	—	—	10	—	—	—
79	Shell Concentration— Type 2	310	80	45	57	—	90	—	—	—	+	—	—	+	—
124	Hearth, Surficial—Type 2	100	100	90	92	—	+	—	—	—	+	—	—	—	—
80	Bone fragment scatter	70	60	127	130		—	—	—	—	—	—	—	—	—
125	Hearths, Multiple	80	80	150	153		—	—	—	—	—	—	—	—	—
16	Shell Concentration— Type 1	70	55	93	99	—	95	—	+	—	—	—	—	+	—
104	Shell Concentration— Type 1	130	80	145	160	—	23	—	—	—	70	5	—	—	+
15	Shell Concentration— Type 1	40	30	143	148	+	—	—	+	—	+	+	—	—	—
81	Hearth, Basin—Type 2	90	20	80	85	—	95	—	—	—	—	—	—	+	—
14	Bone fragment scatter	50	50	125	135	—	—	—	+	—	—	—			+
17	Shell Concentration— Type 2	150	150	85	95	10	80	5		—	—	—	—	—	—
110	Pit—Type 1	100	40	150	165	—	+	—	45	—	—	45	+	+	+
82	Pit—Type 1	100	50	135	140	—	+	—	+	—	—	—	—	—	—
24	Shell Concentration— Type 2	150	120	91	97	—	—	85	—	—	—	15	—	—	—
1	Shell Concentration— Type 1	80	55	75	100	65	+	—	10	—	20	+	—	—	—
11	Shell Concentration— Type 2	300	80	85	110	10	23	+	36	—	+	24	+	—	—
13	Bone bed	170	160	132	150	+	8		67	4	5	14		+	+
25	Shell Concentration— Type 2	300	100	85	94	+	—	—	95	—	—	+	—	—	—
83	Shell Concentration— Type 1	120	80	45	60	—	80	—	—	—	15	—	—	—	—
12	Shell Concentration— Type 2	300	80	90	110	+	+	—	92	—	+	6	—	—	—
10	Shell Concentration— Type 1	150	80	80	100	—	—	—	95	—	+	+	—	—	—
85	Hearth, Surficial— Type 1	37	23	65	67	—	+		—	—	—	—	—	—	—
87	Shell Concentration— Type 1	60	40	100	103	—	—	—	+	+	—	+	—	—	—
26	Shell Concentration— Type 3	600	300	120	138	15	30	—	50	—	—	5	—	—	—
9	Shell Concentration— Type 2	300	60	50	60	17	58	+	17	—	—	—	—	—	+
78	Shell Concentration— Type 1	160	60	71	80	+	+	—	+	+		—	—	—	—

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

* Estimates based on observation in the field. Actual counts from sampling are included in Fauna section of Chapter 11.

“+” sign means presence mentioned in notes; “—” sign means absent/not observed.

charcoal at 124 cmbs, and from OZ-G, were radiocarbon dated. In this section, we only report the dates derived from charcoal or organic sediment. No shell dates are used. Of particular interest among the South Dune features are Features 7, 2, 16, 81, and 10, which are dated, and Feature 13, which is a bone bed.

Feature 7

Feature 7 is an example of the *Rabdotus* cluster feature type (Table 6-3). It is a small cluster of *Rabdotus* land snails with a few broken Lightning Whelks (*Busycon sinistrum*) and Bay Scallops (*Argopectin irradians*), and a scatter of charcoal (Table 6-4). Feature 7 is the southern-most and deepest of the features recorded on the site, and was dated to a time (3470 – 3350 B.P.) when the barrier islands may not have fully developed.

Feature 2

Feature 2 is a thin, linear concentration of Sunray Venus clams (*Macrocallista nimbosa*), most of which are in fragments (Table 6-4). Feature 2 was dated to 1170-960 B.P. (2 Sigma Calibration), making it, along with Feature 123 on the north dune, one of the youngest feature recorded on the site.

Feature 16

Feature 16 was a relatively small concentration of 68 Lightning Whelks with a few other marine shells (Table 6-4) and a small amount of bone, including some from a deer-sized artiodactyls. It was dated to 1300 – 1180 B.P.

Feature 81

Feature 81, only about 15 m north of Feature 16 on the slope of the dune, had a similar date of 1350 – 1270 B.P. Feature 81 was a small, shallow basin hearth with a few scattered shells (Table 6-4).

Feature 10

Feature 10 is a small but dense concentration of oyster shell (*Crassostrea virginica*) with a few other shells (Table 6-4), and large numbers of small bone fragments, including those of fish, a rabbit (*Sylvilagus* sp.), and a rattlesnake (*Crotalus* sp.). It dates from 2710 to 2350 B.P.

Feature 13

This feature was a collection of horse (*Equus caballus*) bones found piled in the bottom of a large erosional gully, filled with laminated layers of sand and very fine dark clay (Figure

6-15a). The bones continued into the wall of the trench. Feature 13 was re-located during Phase III, while realigning Trench B. The bones were, for the most part, disarticulated and in a jumbled pile several bones deep in the narrow bottom of the gully, which extended to 150 cmbs (Figure 6-15b). The bones are those of an adult horse, and represents an MNI of 1. One rib bone has evidence that it was broken and healed completely, with bone smoothly remodeled. There are no butcher marks on the bones. There is no evidence of chewing by either rodents or carnivores. The absence of chewing marks and lack of evidence of atmospheric weathering suggests the bones were buried soon after deposit.

The lowest layer of bones included a humerus and radio-ulna that remained in anatomical position and were probably still articulated when the bone reached the bottom of the gully. The two femurs were not far apart, and in roughly anatomical position, with the right innominate located nearby. The bones were in excellent condition. Eight species of marine shells were identified in close proximity to the horse bones (Figure 6-15b), including an unusual number (n=11) of banded tulips (*Fasciolaria lillium*), a species that is only occasionally found in features at the site. Also present were a number of Lightning Whelks, and more than 300 oyster shell fragments, representing at least 41 individuals. The presence of so many shells, as well as the charcoal and burned clay, may be the result of deflation. However, the presence of at least a few articulated bones (Figure 6-15c), the uncommon shell type, and the generally good condition of the bone suggests that the contents could not have suffered significant movement after deposition. It is possible that a horse was buried soon after its death.

Features on the Beaumont Ridge

The south dune is separated from the north dune by a remnant ridge of Beaumont Clay. When undisturbed, this Pleistocene clay has a layer of cumulic sandy clay above it (Ricklis 1997:17). Within this cumulic clay, usually sitting directly above the Beaumont clay, 19 features were encountered during the project (Table 6-5). Feature locations are shown on Figures 6-6 and 6-7. Of particular interest are Features 28, 31, 33/121, 53, 54, 57 and 122. Features 31, 54, 33/121, and 122 have associated charcoal dates. The features are discussed below as they occurred, south to north.

Feature 57

Feature 57, the largest feature in the Beaumont Ridge area, was a moderately thick concentration composed mostly of oysters, with an area on the northern end that was almost entirely scallops (Figure 6-16; Table 6-5).

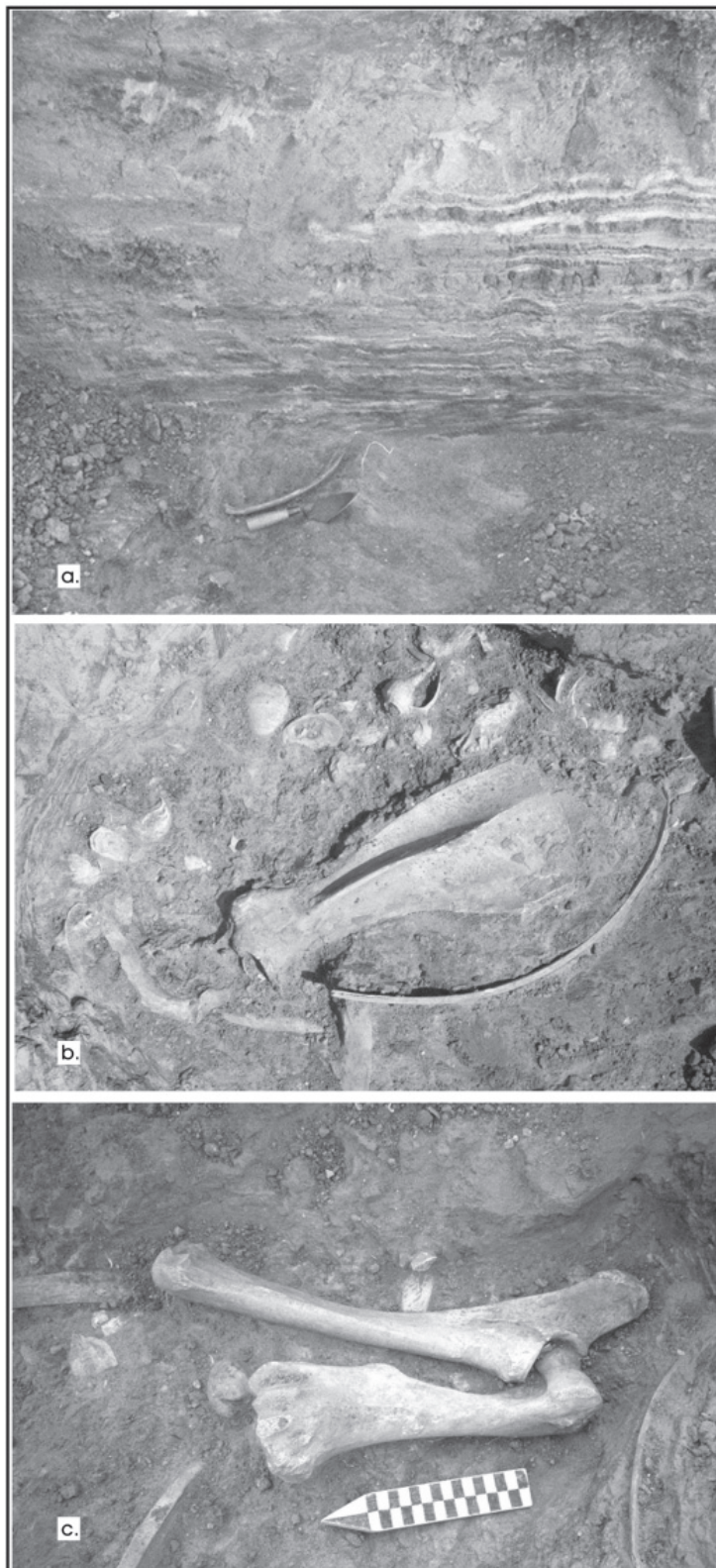


Figure 6-15. Feature 13: a) profile showing the naturally filled gully in which Feature 13 was located; b) after clearing of the upper layer of horse bones; c) the lower level of horse bones. Note that the humerus and radio-ulna are in anatomical position.

Feature 54

Feature 54 was one of the more interesting features at the site. First seen as a dark stain in medium light clay, truncated on the northwest by the digging of the western half of Trench B (Figure 6-17a), it was found to be a large, complex basin hearth, dug into the Beaumont clay (Figure 6-17b). The south half had a layer of charcoal lining the basin that was at least 2 cm thick and was filled with lumps of burned clay and charcoal. The north basin appeared to have had a similar charcoal lining at one time, but it had been dug out, leaving layers of ash, charcoal, and a few Lightning Whelk shells (Table 6-5). The charcoal in this feature dated to 1290 – 1160 B.P.

Feature 28

Feature 28 was not immediately noticed among the scattered shells just above the Beaumont, but was conveniently cut in half by the Gradall, providing an excellent profile. It was dug more than 15 cm into the Beaumont clay and filled with oyster shell (Table 6-5). Besides oysters there were only a few other marine shell fragments, a few flecks of burned clay, and charcoal in the pit fill.

Feature 31

Feature 31 was a fairly large hearth (Table 6-5) with a shallow (ca. 6 cm) basin. It was dug from the cumulic sandy clay almost to the top of the Beaumont clay. The basin looks as if it had been scooped out at least once (Figure 6-18). When abandoned, it was left packed with charcoal and small (ca. 1-2 cm), soft burned clay nodules. This feature was dated to 1820 – 1600 B.P.

Feature 33/121

Feature 33 and 121, defined separately in the field were actually the same feature, and will be considered together here and in Table 6-5. In Trench B, Feature 33 appeared to be a pile of Lightning Whelk shells. As it extended to the east, however in Trench I, this western pile was seen to be the largest of three whelk shell piles surrounded by a thinner scatter of oysters and quahogs (Table 6-5). The eastern part of this feature was dated to 2730 – 2360 B.P.

Feature 53

Feature 53 was a fairly large pit with flat bottom and uneven walls, 10 cm deep and completely packed with shell. Most of these shells are large *Rabdotus* land

Table 6-5. Features on the Beaumont Ridge

Feature	Feature Type	Description				Estimated % of Common Shells*									
		Length (cm)†	Width (cm)†	Depth at top (cm)‡	Depth at Bottom (cm)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria tulium</i>	Sunray Venus Clam <i>Macrocallista nimbosa</i>	Southern Quahog <i>Mercenaria campechianensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.
86	Shell Concentration— Type 3	400	100	37	53	5	—	—	20	—	—	75	—	—	—
88	Shell Concentration— Type 1	78	65	92	110	—	—	—	50	—	—	45	—	—	—
56	Hearth, Basin— Type 1	30	30	46	60	—	—	—	—	—	—	—	—	—	—
57	Hearth Basin —Type 1	500	200	60	68	20	5	—	70	—	—	—	—	—	—
27	Hearth, Surficial— Type 1	40	40	65	80	—	+	—	+	—	—	—	—	—	—
89	Shell Concentration— Type 1	75	42	67	69	—	—	—	—	—	—	—	—	—	—
130	Shell Concentration— Type 3	260	200	50	60	—	+	—	—	—	+	+	—	—	—
55	Hearths, Multiple	140	120	43	45	—	—	—	—	—	—	—	—	—	—
54	Hearth, Basin—Type 2	160	75	53	80	—	+	—	—	—	—	—	—	—	—
28	Pit—Type 1	58	28	55	70	10	—	—	90	—	—	—	—	—	—
91	Hearth, Basin—Type 2	95	95	75	108	—	—	—	—	—	—	—	—	—	—
144	Hearth, Basin—Type 2	380	300	25	28	—	+	—	—	—	70	20	—	—	—
29	Hearth, Basin—Type 1	40	30	72	89.3	—	+	—	+	—	—	—	—	—	—
31	Hearth, Basin—Type 1	50	50	84	90	—	+	—	+	—	—	—	—	—	—
32	Hearth, Basin—Type 1	150	150	70	78	—	30	—	60	—	—	—	—	—	—
33/121	Shell Concentration— Type 2	200	140	105	125	+	48	—	25	—	—	25	—	—	—
53	Pit—Type 3	60	40	113	124	+	15	—	10	+	—	—	—	—	70
122	Pit—Type 1	80	30	105	125	—	95	—	4	—	—	—	—	—	+
52	Shell Concentration— Type 3	500	200	56	60	—	60	—	10	—	—	29	+	—	—

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

* Estimates based on observation in the field. Actual counts from sampling are included in Fauna section of Chapter 11.

“+” sign means presence mentioned in notes; “—” sign means absent/not observed.



Figure 6-16. The numerous scallops in the north end of Feature 57, in the Beaumont Ridge area. Note the numerous burned clay nodules.

snails, but fragments of marine shells were also present (see Table 6-5). Charcoal and small (ca. 1 cm) burned clay nodules were present in the fill.

Feature 122

Feature 122 is an example of a pit packed with shells that were also piled above and around the pit. In this case, the shells were Lightning Whelk, with a scatter of *Rabdotus* and a few oysters (Table 6-5). The feature was dated to 2320 – 2060 B.P.

Features in the North Dune

Table 6-6 presents basic data on the 103 features in the north dune, as well as on 2 features (147, 162) in the pond area. Locations of features are shown on Figures 6-5 and 6-6. The discussion is organized by location,



Figure 6-17. Feature 54, a large hearth in the Beaumont Ridge area: a) the top of Feature 54 (note western half of Trench B, visible at the top of the picture, had reached the Beaumont, truncating the northern corner of the feature); b) a profile showing the double basin and the layer of charcoal in the southern basin that was partially removed in the northern basin.



Figure 6-18. Profile of Feature 31, showing hearth filled with burned clay nodules and charcoal.

Table 6-6. Features in North Dune and the Pond/ Northwest Ridge Area

Feature	Feature Type	Description				Estimated % of Common Shells*									
		Length (cm)†	Width (cm)†	Depth at top (cmbs)‡	Depth at Bottom (cmbs)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria litum</i>	Sunray Venus Clam <i>Macrocallista nimbosa</i>	Southern Quahog <i>Mercenaria campechensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.
123	Shell Concentration— Type 2	140	120	65	85	—	90	—	—	—	5	5	—	—	—
76	Cluster of sunray shell chipping	150	150	95	100	—	—	—	—	—	+	—	—	—	—
34	Hearth, Surficial— Type 1	30	30	74	76	—	—	—	—	—	—	—	—	—	—
35	Hearth, Basin—Type 1	50	50	98	124	—	—	—	—	—	—	—	—	—	—
36	Shell Concentration— Type 1	70	58	59	62	—	90	—	—	—	—	—	—	+	+
37	Hearth, Surficial— Type 2	75	75	56	60	—	90	—	+	—	—	—	—	+	+
38	Shell Concentration— Type 3	220	220	113	118	—	15	—	80	—	—	+	—	+	+
157	Shell Concentration— Type 1	100	100	150	153	—	+	—	95	—	—	—	—	—	—

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

* Estimates based on observation in the field. Actual counts from sampling are included in Fauna section of Chapter 11.

“+” sign means presence mentioned in notes; “—” sign means absent/not observed.

Table 6-6. Continued...

Feature	Feature Type	Description				Estimated % of Common Shells*									
		Length (cm)†	Width (cm)†	Depth at top (cmbs)‡	Depth at Bottom (cmbs)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria litium</i>	Sunray Venus Clam <i>Macrocallysta nimbosa</i>	Southern Quahog <i>Mercenaria campechiensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.
90	Shell cluster	—	—	90	110	—	+	—	95	—	—	+	+	+	+
101	Hearth, Basin— Type 2	70	70	78	83	—	+	—	—	—	—	—	—	—	—
39	Hearth, Surficial— Type 1	60	50	87	88	—	—	—	—	—	—	—	—	—	—
137	Hearth, Basin— Type 2	500	265	138	145	—	—	—	—	—	—	—	—	—	—
41	Pit—Type 3	30	30	74	78	—	30	50	15	—	—	—	—	—	+
148	Pit—Type 1	40	40	78	84	—	35	—	60	—	—	+	—	+	—
42	Hearth, Surficial— Type 1	50	28	74	76	—	—	—	—	—	—	—	—	—	—
43	Hearth, Surficial— Type 1	30	30	72	74	—	+	—	—	—	—	—	—	—	—
44	Hearth, Surficial— Type 1	60	50	78	83	—	+	—	+	—	—	—	—	—	—
102	Pit—Type 1	100	75	61	71	—	85	—	—	—	10	—	—	—	5
103	Shell Concentration— Type 2	200	100	80	82	—	80	—	10	—	10	—	—	—	—
45	Hearth, Surficial— Type 1	30	20	52	55	—	—	—	—	—	—	—	—	—	—
46	Hearth, Basin—Type 1	30	30	70	75	—	+	—	—	—	—	—	—	—	—
141	Shell Concentration— Type 2	350	50	95	105	—	5	—	—	—	90	—	5	—	—
49	Hearth, Basin—Type 1	37	32	58	62	—	+	—	+	—	—	—	—	—	—
47	Shell Concentration— Type 1	94	65	60	62	—	—	—	+	—	—	—	—	—	—
48	Hearth, Basin—Type 4	79	70	58	75	—	—	—	—	—	—	—	—	—	—
50	Hearth, Basin—Type 1	60	57	64	66	—	—	—	—	—	—	—	—	—	—
51	Pit—Type 3	20	10	59	63	—	—	—	—	—	—	—	—	—	+
142	Shell Concentration— Type 3	410	300	15	31	5	85	—	—	—	—	5	—	—	5
58	Hearth, Basin— Type 2	40	20	67	73	—	—	—	—	—	+	—	—	—	—
107	Rabdotus cluster	120	40	65	68	—	—	—	—	—	—	—	—	—	85
59	Hearth, Basin— Type 1	90	80	93	99	—	+	—	—	—	+	—	—	—	—
105	Hearth, Surficial— Type 1	40	40	60	62	—	—	—	—	—	—	—	—	—	—
106	Hearth, Basin— Type 2	100	80	95	102	—	—	—	—	—	—	—	—	—	—
98	Hearth, Surficial— Type 2	60	45	40	45	—	+	—	—	—	+	+	+	—	—
60	Hearth, Basin— Type 1	60	40	210	215	—	—	—	+	—	—	—	—	—	—
96	Bone fragment scatter	550	200	40	43	—	—	—	—	—	—	—	—	—	—
95	Shell Concentration— Type 3	400	150	60	70	—	50	—	+	+	40	8	—	—	—
61	Burial pit	175	75	54	73	—	—	—	—	—	+	—	—	—	—
161	Shell Concentration— Type 1	160	60	25	30	—	20	—	80	—	—	—	—	—	—
97	Hearth, Surficial— Type 2	100	100	42	45	—	80	—	—	—	20	—	—	—	—
151	Shell Concentration— Type 1	150	50	57	62	—	—	—	—	—	—	—	—	—	40

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

* Estimates based on observation in the field. Actual counts from sampling are included in Fauna section of Chapter 11.

“+” sign means presence mentioned in notes; “—” sign means absent/not observed.

Table 6-6. Continued...

Feature	Feature Type	Description				Estimated % of Common Shells*									
		Length (cm)†	Width (cm)†	Depth at top (cmbs)‡	Depth at Bottom (cmbs)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria lilium</i>	Sunray Venus Clam <i>Macracallista nimbosa</i>	Southern Quahog <i>Mercenaria campechiensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.
63	Charcoal cluster	108	80	57	60	—	—	—	—	—	+	—	—	—	+
152	Shell Concentration— Type 1	40	30	160	168	—	—	—	100	—	—	—	—	—	—
100	Shell Concentration— Type 1	100	20	42	48	—	+	—	—	—	25	—	—	—	70
64	Hearth, Basin— Type 1	70	40	65	80	—	—	—	—	—	—	—	—	—	—
153	Shell Concentration— Type 1	170	5	44	50	—	—	—	+	—	95	—	—	—	—
154	Hearth, Surficial— Type 2	70	50	46	52	—	50	—	—	—	50	—	—	—	—
67	Hearth, Basin— Type 1	43	37	120	130	—	—	—	—	—	—	—	—	—	—
65	Shell Concentration— Type 1	50	40	50	52	+	+	—	+	+	+	—	—	—	+
69	Pit—Type 4	140	80	50	240	—	—	—	—	—	—	—	—	—	—
68	Hearth, Basin—Type 1	40	?	112	119	—	—	—	—	—	—	—	—	—	—
70	Charcoal cluster	60	45	60	62	—	—	—	—	—	+	—	+	—	—
71	Hearth, Surficial— Type 1	30	30	144	147	—	—	—	—	—	—	—	—	—	—
72	Hearths, Multiple	200	200	119	123	—	+	—	—	—	+	—	—	—	—
73	Pit—Type 3	33	27	45	48	—	—	—	—	—	—	—	—	—	90
158	Shell Concentration— Type 1	30	20	180	183	—	—	—	100	—	—	—	—	—	—
156	Hearth, Basin— Type 2	70	50	42	55	—	20	—	50	—	—	10	—	—	20
75	Pit—Type 1	67	?	57	90	—	+	—	+	—	+	—	+	—	—
108	Hearth, Surficial— Type 2	300	70	35	39	—	80	—	—	—	10	—	—	—	10
74	Pit—Type 1	45	30	60	63	—	+	—	—	—	—	—	—	—	+
109	Shell Concentration— Type 2	350	80	70	90	—	80	—	5	—	15	—	—	—	—
77	Shell Concentration— Type 1	30	30	220	223	—	—	—	—	—	—	—	—	—	—
92	Shell Concentration— Type 2	140	140	240	260	60	+	—	35	—	—	—	—	—	—
159	Pit—Type 1	150	100	30	50	—	10	—	45	—	45	—	—	—	—
113	Hearth, Surficial— Type 2	60	30	98	104	—	90	—	—	—	10	—	—	—	—
112	Hearth, Surficial— Type 2	20	20	117	118	—	—	—	—	—	—	—	—	—	—
111	Pit—Type 1	70	45	80	85	—	90	—	—	—	10	—	—	—	—
143	Shell Concentration— Type 1	50	?	121	130	—	+	—	—	—	—	95	—	—	+
118	Hearth, Basin— Type 1	20	20	50	53	—	+	—	—	—	—	—	—	—	—
119	Hearth, Basin— Type 1	35	25	70	77	—	—	—	—	—	—	—	—	—	—
120	Shell Concentration— Type 1	100	50	68	74	—	37	—	+	—	60	—	—	—	—
114	Hearths, Multiple	650	100	18	29	—	—	—	—	—	—	—	—	—	—
117	Pit—Type 1	130	50	32	42	+	+	—	—	—	—	—	—	—	+
115	Hearth, Basin— Type 1	70	40	70	72	—	—	—	—	—	—	—	—	—	—

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

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“+” sign means presence mentioned in notes; “—” sign means absent/not observed.

Table 6-6. Continued...

Feature	Feature Type	Description				Estimated % of Common Shells*										
		Length (cm)†	Width (cm)†	Depth at top (cmbs)‡	Depth at Bottom (cmbs)‡	Atlantic Bay Scallop <i>Argopectin irradians</i>	Lightning Whelk <i>Busycon sinistrum</i>	Crossbar Venus Clam <i>Chione cancellata</i>	Eastern Oyster <i>Crassostrea virginica</i>	Banded Tulip <i>Fasciolaria tulium</i>	Sunray Venus Clam <i>Macrocallysta nimbosa</i>	Southern Quahog <i>Merrenaria campechiensis</i>	Florida Horse Conch <i>Pleuroploca gigantea</i>	Shark's Eye <i>Polinices duplicatus</i>	Rabdotus land snail <i>Rabdotus</i> sp.	
116	Hearth, Surficial— Type 2	100	60	105	107	—	—	—	—	—	—	—	—	—	+	
126	Shell Concentration— Type 2	200	150	50	52	—	—	—	—	—	—	—	—	—	—	
93	Shell Concentration— Type 1	140	80	50	53	—	24	—	24	—	24	24	—	—	+	
94	Hearth, Basin— Type 2	350	200	90	120	—	12	—	50	—	35	—	—	—	+	
99	Shell Concentration— Type 1	60	23	69	79	—	35	—	55	—	10	—	—	—	—	
131	Hearth, Basin— Type 1	25	20	94	96	—	—	—	—	—	—	—	—	—	—	
132	Shell Concentration— Type 3	220	200	46	63	—	+	—	—	—	—	—	—	+	—	
139	Shell Concentration— Type 3	250	150	55	65	—	+	—	+	—	+	—	—	—	—	
134	Shell Concentration— Type 2	250	100	40	55	—	50	—	50	—	—	—	—	—	—	
136	Old backhoe trench containing disarticulated human bones (Burial #4)	220	50	30	110	—	—	—	—	—	—	—	—	—	—	
140	Charcoal cluster	150	75	72	73	—	—	—	—	—	—	—	—	—	+	
146	Hearth, Basin—Type 2	80	50	60	64	—	—	—	—	—	—	—	—	—	—	
138	Burial pit containing Burials 6 and 7	60	45	80	85	—	—	—	—	—	—	—	—	—	—	
133	Hearth, Surficial— Type 2	800	150	65	68	—	5	—	—	—	95	—	—	—	—	
145	Bone fragment scatter	150	100	70	73	—	—	—	—	—	—	—	—	—	—	
149	Shell Concentration— Type 3	500	100	52	55	—	30	—	70	—	—	—	—	—	—	
66	Hearth, Surficial— Type 1	58	38	60	62	—	—	—	—	—	—	—	—	—	—	
155	Pit—Type 1	70	40	46	50	—	10	—	50	—	—	—	—	—	40	
40	Hearth, Surficial— Type 2	30	20	83	86	—	—	—	+	—	—	—	—	—	—	
150	Shell Concentration— Type 1	75	60	45	60	—	+	—	95	—	—	—	—	—	—	
160	Shell Concentration— Type 1	50	40	45	47	—	70	—	—	—	—	30	—	—	—	
147	Pit, Multiple	200	100	15	60	—	30	—	60	—	+	+	+	—	—	
163	Pit—Type 1	40	20	30	40	—	60	—	25	—	—	10	5	—	—	
162	Shell Concentration— Type 1	75	75	70	72	—	70	—	30	—	—	—	—	—	—	
18	Shell Concentration— Type 3	300	300	70	80	5	15	—	—	—	80	—	—	—	—	
165	Shell Concentration— Type 2	300	90	35	40	—	10	—	—	+	85	—	—	—	—	
164	Bone fragment scatter	180	90	40	50	—	—	—	—	—	—	—	—	—	—	
166	Hearth, Basin— Type 2	100	100	50	53	—	—	—	—	—	—	—	—	—	—	
167	Shell Concentration— Type 2	200	100	50	56	—	+	—	+	—	—	—	—	—	+	
169	Shell Concentration— Type 1	120	100	140	145	—	—	—	—	—	—	—	—	—	+	
170	Bone bed	80	70	85	110	—	—	—	—	—	—	—	—	—	—	

† Measurements within the boundaries of the trench.

‡ Vertical datum in the field was the road or ground surface.

* Estimates based on observation in the field. Actual counts from sampling are included in Fauna section of Chapter 11.

“+” sign means presence mentioned in notes; “—” sign means absent/not observed.

south to north, and includes an in-field estimate of the kinds of shells found in each. Artifacts located in these features were discussed in the previous chapter. Analysis of the faunal remains will be included in Chapter 11. Of particular interest are Features 18, 41, 46, 49, 63, 64, 72, 75, 76, 103, 108, 109, 113, 115, 116, 123, 137, 153, 155, and 170.

Feature 123

Feature 123 was located at the southern end of the north dune and consisted of a 10 cm thick layer of Lightning Whelks (Table 6-6), charcoal, and burned clay nodules. Charcoal from this feature dated to 1170-960 B.P.

Feature 76

Feature 76 consisted of a small scatter of charcoal, one broken Sunray shell tool (see Artifacts Section, Chapter 5) and about 8-10 thin flakes of Sunray shell, broken in a way that was not commonly seen. In the field it was considered possible that they were the debitage from manufacture of chipped shell tools. Examination of these shell fragments in the lab, under a microscope did not provide proof that they were, indeed, shell debitage, but the fact remains that Sunray shells normally break perpendicular to the thickness of the shell, while these chips were broken longitudinally along the thickness of the shell. It remains possible, therefore that these shell fragments were deliberately struck from Sunray shells in the process of shell modification.

Feature 137

Feature 137 was a small basin hearth surrounded by animal bone and a few shell fragments (Table 6-6). Most of the animal bone appeared to be deer-sized, but it was highly fragmented. The features dated to 2730-2350 B.P.

Feature 41

Feature 41 is an example of Pit Type 3, a shallow, flat-bottomed pit filled with charcoal and shells, but this feature was unique. The feature was roughly 30 cm x 30 cm, about 5-6 cm deep, and had a roughly square shape (Figure 6-19). Twenty four species with an MNI of 98 were identified in this small space (Table 6-6). Most were small species, many of which were only found in this feature. Even the examples of larger species were generally very young and small. The two species that were by far the most common were oysters, many of which were no bigger than "spats", the size at which oysters cease floating and seek out a permanent location, and Crossbar Venus (*Chione cancellata*). The later is a common species, but is generally no bigger than 4.4 cm (1.75"), and is

thus not considered a significant food resource. Details about the species richness in this feature are discussed in Chapter 11.

Feature 103

Feature 103 was a distinct but thin shell concentration, consisting mostly of Lightning Whelks, with a few oysters and Sunray Venus shells (Table 6-6). Charcoal from this feature dated to 2300-1900 B.P.

Feature 46

Feature 46 was a typical shallow basin hearth with a few shells scattered nearby. It was dated to 2050-1880 B.P.

Feature 49

Feature 49 is also a shallow basin hearth, much smaller than Feature 46. Its 3-4 cm basin is filled with charcoal and a few shells. It dates to 1870-1700 B.P.

Feature 63

Feature 63 is an example of a charcoal scatter that does not appear to be a hearth. There was a dense patch of charcoal and a few shells surrounded by a thinner scatter of charcoal. This feature was dated to 3070-2800, and is the oldest dated feature in the North Dune.

Feature 64

Feature 64 is perhaps the best example of a basin hearth filled with numerous burned clay nodules and charcoal (Figure 6-20). The nodules range in size from 2-7 cm in maximum dimension.

Feature 153

Feature 153 was located on the west side of Ennis Joslin Road, in a paleosol above the dark layer containing OZ-J. It was a dense layer of Sunray Venus clam shells, all laying interior-down and piled several deep (Table 6-6). This feature was dated to 2300-2000 B.P.

Feature 72

Feature 72 is an unusual feature because it was found in a light clay layer, while most cultural manifestations at 41NU2 were found in dark paleosols. The feature consists of 2 surficial hearths and 2 charcoal concentrations, all



Figure 6-19. Feature 41: a) when first located; b) after removal of contents. Note the unusually square shape.

within 2 m of each other and surrounded by a scatter of unusually large lumps of charcoal.

Feature 75

This pit was dug from the center level of OZ-D, but the original shape cannot be determined because the upper part of the pit was not noticed among the general shell and charcoal scatter typical of this occupation zone. The pit was identified when the digging reached the light clay layer, where several large shells were seen (Figure 6-21a). The walls of the pit were then visible in the trench wall (Figure 6-21b). The upper part is mostly charcoal, burned clay, animal bones (including several deer bones), with some fragmented shell. Lower part is crammed with whole shells, an extremely large Lightning Whelk, and many other species (Table 6-6).

Feature 108

Feature 108 was small surficial hearth with a dense scatter of shell, consisting mostly of Lightning Whelks, although some parts of the feature had large numbers of *Rabdodus* (Table 6-6). This feature was in a paleosol that had been removed by

road construction in most of Trench B, and lays above OZ-D. It was dated to 1410-1300 B.P.

Feature 109

Feature 109 was a 10-cm thick pile of Lightning Whelks, with a few oyster and Sunray shells laying directly on a slight mound of light-colored clay (Table 6-6). This feature was an unusually thick and dense shell accumulation in which most of the shells were abnormally large.

Feature 113

Feature 113 was a small surficial hearth surrounded by Lightning Whelks and a large number of fish and mammal bones. Two large sea catfish otoliths were found side-by-side in this feature. It was dated to 2130-1930 B.P.

Feature 115

Feature 115 is a small (ca. 20 cm in diameter) hearth surrounded by a thin scatter of charcoal and shell, in a relatively light clay layer. It dated to 2770-2720 B.P.



Figure 6-20. Feature 64, showing close-up of hearth basin filled with burned clay nodules.



Figure 6-21. Feature 75: a) when first observed; b) east wall of trench showing the profile of the pit, dug from OZ D.

Feature 116

Feature 116 was located 30 cm below Feature 115, in a light clay layer. It consists of a small hearth, with a scatter of charcoal, bone, *Rabdotus* and Sunray Venus fragments. In many ways, it is reminiscent of Feature 72 in that it was found in a light clay layer and the charcoal fragments are unusually large. The lowest part of the feature contained large bone fragments, large charcoal pieces and a cluster of *Rabdotus* shells (Table 6-6).

Feature 155

Feature 155 was a small, dense concentration of oysters, *Rabdotus*, and Lightning Whelk shells (Table 6-6). The feature dated to 2480-2340 B.P.

Feature 170

As mentioned in the narrative for Trench T in the previous chapter, Feature 170 was discovered when the backhoe operator dug a large scoop of sediment from beneath a curb in the parking lot north of Ocean Drive. This was not

the usual procedure, but the operator had been requested to avoid breaking the curb in this location. In order to dig beneath it, he was forced to dig out a larger-than-normal scoop. When this bucket was emptied, numerous large bones were observed. These were collected, and the area under the curb was examined. The bone bed was found to consist of the lower legs of bovids, most of which were still articulated. The bone bed extended into the east wall of the trench.

Most of the articulated segments ended at the proximal end of the metapodials, though a few carpals and tarsals were also present. Table 6-7 is a complete list of bones recovered within Trench T. All bones identifiable to the Genus taxonomic level were *Bos taurus*. The bones were in a sandy matrix unlike the clay dune surrounding the area. Further excavation the next morning exposed more articulated bone (Figure 6-22), as well as several cranium fragments, and three horn cores, including one that was 10 cm in diameter and extended 45 cm within the trench and continued into the eastern wall. This horn made it likely that the bones were cattle (*Bos taurus*), as bison (*Bison bison*) horns are never that long. Unfortunately, all horn cores disintegrated when they were removed.

Table 6-7. Bovid Elements Identified from Feature 170

Element	L	R	Total	Notes
ELEMENTS AT OR NEAR MATURE SIZE, MNI = 6				
Maxilla (fragment)	1	—	1	—
2+3 Carpal	—	1	1	—
Auxillary Carpal	—	1	1	—
Metacarpal	5	5	10	One has unsealed distal epiphysis, MNI =6;
Central +4th Tarsal	1	—	1	Chopped with large, sharp metal tool;
2+3 Tarsal	2	—	2	—
Metatarsal	2	5	7	One is unsealed and very small;
1st Phalange	—	—	29	Two are unsealed (<18 mo. old);
2nd Phalange	—	—	27	Three are partially sealed (ca. 18 mo. Old);
3rd Phalange	—	—	16	—
Proximal Sesmoids	—	—	20	—
Distal Sesmoids	—	—	8	—
Fragments	—	—	8	—
Total	11	12	131	—
PERINATAL ELEMENTS, MNI = 1				
Radial Carpal	1	—	1	—
Metatarsal	1	—	1	Very small, central diaphysis very thin, distal epiphyses unsealed;
Metapodial Epiphyses	—	—	3	—
1st Phalange	—	—	4	—
2nd Phalange	—	—	6	MNI=1
3rd Phalange	—	—	6	—
Proximal Sesmoids	—	—	1	—
Distal Sesmoids	—	—	1	—
Fragments	—	—	9	—
Total	2	—	32	—
FETAL ELEMENTS, MNI = 1				
Cranium	—	—	5	—
Humerus	1	—	1	Proximal epiphysis;
Radius (fragment)	1	—	1	—
Lumbar Vertebra	—	—	1	Unsealed epiphysis of centrum;
Femur	1	1	2	—
Tibia	1	1	2	—
Ischium	1	—	1	—
1st Phalange	—	—	1	—
2nd Phalange	—	—	1	—
3rd Phalange	—	—	3	—
Fragments	—	—	19	—
Total	5	2	37	—
Overall total	18	14	200	—
Total MNI	—	—	8	—



Figure 6-22. *Feature 170: a) wall of gully after removal of bone within Trench T. Note laminated sand and clay fill and white sand beneath bone bed.*

Included in the bone bed were the lower leg bones of at least two juvenile animals, one of which was very small and probably of perinatal age, and the other was closer to adult size. Three 2nd phalanges were partially sealed, indicating an age of about 18 months for at least one individual. A complete fetus was also discovered. Though much too small to identify its species, it was also likely cattle. The fetus was in fragile condition, due to the incomplete ossification of its skeleton, and only a few of the larger long bones, a few phalanges, and a few cranial fragments were recovered partially intact (Table 6-7).

Further clearing of the bone bed made it clear that the bones lay in a pile at the bottom of a gully that had eroded the edge of the dune (see Figure 6-22). The sand matrix was laminated with thin layers of clay, indicating that the gully had filled in naturally and that the bone was not deliberately buried at the time of disposal. This was confirmed by evidence of extensive atmospheric weathering on one side of some of the bone. The bone bed continued into the east

wall of the trench and remains beneath the driveway of the restaurant north of Ocean Drive.

Feature 18

Feature 18 was a thin scatter of broken shell and charcoal surrounding a dense concentration consisting mostly of Sunray Venus (*Macrocallista nimbosa*) clam shells, all laying interior down and stacked as much as 5 or 6 deep. The Sunray shells are unusually large, with many 10- to 12 cm long. Mixed with them are a few Lightning Whelks and Bay Scallops (Table 6-6). This feature was dated to 1940-1740 B.P.

Features in the Pond and Northwest Ridge Area

Two features in the northern part of the site were not in a clay dune setting. Feature 147 was on the slope of the Beaumont ridge in the northwest part of the site, and Feature

162 was located immediately adjacent to the west of the pond sediments.

Feature 147

Feature 147 was located in the dark cumulic sediment just above the Beaumont clay, situated west of the ephemeral pond. This was not in the north dune, which did not extend west of the pond in this area. Feature 147 consisted of a series of 4 shallow pits filled with oyster and Lightning Whelk shell and surrounded by a thick layer of shell (Table 6-6). Three of the pits were 5 to 7 cm deep, stopping at the Beaumont clay, but the fourth was 15 cm deep and was dug into the light Beaumont layer.

Feature 162

Feature 162 was an unusually dense collection of oyster and Lightning Whelk shell located with OZ-K on the western edge of the ephemeral pond in the northern part of the site (Table 6-6). The feature was badly impacted by the original construction of Ocean Drive, and by the scraping of the new free right turn lane on to Ennis Joslin being constructed.

Occupation Zones

During the digging of Trench B north of about 34+000 and roughly 70 cm below the road surface, monitors noticed a continuous, thin layer of scattered shell, small flecks of charcoal, burned clay, and animal bone in one of the dark clay paleosols. Within this layer were a number of features, including hearths, pits filled with shell, and areas of dense shell accumulation, each surrounded by the thinner, more generalized scatter of cultural remains. Although the dark clay paleosol extended farther north, the cultural deposit ended at about 34+042. In later excavations, similar deposits were seen, especially farther north, where as many as three strata containing such sheet midden deposits were observed in a single profile. These areas of cultural deposits were distinctive, always seen in dark sediment layers, and seemed to represent areas of general occupation refuse within which distinctive, well-defined features could sometimes be identified. These areas often extended more than 10 m within the trench, and although the density of deposits varied over that distance, they were usually clearly definable both vertically and horizontally. These areas were eventually designated Occupation Zones (OZ) and were given letter designations. In all, 11 such zones were identified on the project. The locations of the occupation zones are shown on Figures 6-5, 6-6, and 6-8.

Occupation Zone A

Location: Trench C
Extent: 33+558 to 33+576
Depth: 70-80 cm
Feature(s): 17
Estimated Date: Unknown

OZ-A was defined retrospectively, long after it was observed. Field notes described in the second interim report (Meissner 2003) took on a new meaning in the light of the Occupation Zones that had been defined while Trench B was excavated. These notes recorded a paleosol in Trench C that contained a thin scatter of shell, charcoal, burned clay flecks and numerous highly fragmented animal bones. This scatter remained consistent for 18 m in Trench C. Although the density of cultural remains varied, it was never defined enough to consider it a feature. This description clearly matched the definition of Occupation Zone developed while digging Trench B. The OZ in Trench C was therefore designated OZ-A. Within this thin scatter one feature, Feature 17, was defined.

Occupation Zone B

Location: Trenches B, I
Extent: 34+000 to 34+042
Depth at Top: 70-80 cm
Feature(s): 40—43, 45, 46 and 103
Estimated Date: 2300-1880 B.P.

OZ-B was a layer about 10-12 cm thick of scattered shell, small fragments of fish and deer-sized mammal bone, charcoal flecks, and burned clay flecks extending 42 m in Trench B in a dark clay paleosol that also extended east into Trench I in the southern-most part of the zone. The degree of shell concentration varied over the length of the zone, but was never dense (Figure 6-23). As discussed previously, a dart point that is probably a heavily re-sharpened Matamoros point was found in OZ-B near its southern end. The surface of the paleosol dropped slowly to the north, so that the top was at 70 cm brs in the southern end and 80 cm brs at the northern end in Trench B and somewhat lower in Trench I.

The features in OZ-B included seven small hearths, three of which were built directly on the ground surface with no pit, while two had a 5 cm basin. Feature 40 was a shallow pit, dug from OZ-B and filled with numerous small shells and charcoal. Charcoal from Feature 46 dates OZ-B to 2050 – 1880 B.P. (2 Sigma Cal.), while organic sediment from Feature 103 was dated slightly older at 2300 – 1900 B.P.



Figure 6-23. Scattered shell, charcoal and bone fragments in OZ-B.

Trench I, it was truncated on the west by the setting trench for the curb on the eastern edge of Alameda. The zone contained Features 98 and 100. It is suspected that portions of OZ-C did at one time extend farther to the west, but were truncated by construction of Alameda Street. It is possible that Feature 73, seen in Trench B located in a dark sediment sometimes seen above OZ-D, was part of OZ-C before the building of Ennis Joslin.

Occupation Zone D

Location: Trenches B, I
 Extent: 34+078 to 34+103
 Depth: 50-60 cm
 Feature(s): 65, 69 (?), and 75
 Estimated Date: 1930-1730 B.P.

This occupation zone varied in density more than most of the zones identified. It covered nearly 40 m and consisted of a layer of shell, small fragments of animal bone, charcoal and burned clay. It often seemed to be ending, then would become denser again. It contained Features 65 and 75. Feature 69 may or may not be associated with OZ-D. It was an almost 2-m-wide pit that appeared to have been dug from the top of the paleosol containing OZ-D. However, it is possible that it was dug through the lighter sediments above OZ-D and was modern in origin. Organic sediments taken from OZ-D were dated 1930-1730 B.P. (2 Sigma Cal.).

Occupation Zone C

Location: Trench I
 Extent: 34+045.5 to 34+073
 Depth: 70-80 cm
 Feature(s): 98, 100
 Estimated Date: Unknown

This zone consisted of a dense layer of shell, animal bone fragments, and charcoal. The western edge of the zone was often clearly seen in the trench (Figure 6-24). The eastern edge (within the trench) was truncated by an old gas pipe trench, and when the zone did extend completely across

Occupation Zone E

Location: Trenches B, I
 Extent: 34+080.3 to 34+087
 Depth: 60-65 cm
 Feature(s): 70
 Estimated Date: 1980-1820 B.P.

OZ-E was relatively small for an OZ, but otherwise was similar to OZ-D just above it. Only one feature, Feature 70, was defined within this OZ. OZ-D and -E were separated by a light sediment layer. While OZ-E seemed to lay on a relatively flat sediments, it is clear



Figure 6-24. OZ-C in Trench I, showing density of shells and sharply delineated western edge.

that the light sediments that built up between OZ-D and -E were deeper to the west, so that OZ-D developed on more sloped sediments. To the east, in Trench I, it was sometime difficult, given the method of excavation, to separate them.

Organic sediments in OZ-D and charcoal from OZ-E, respectively, yielded dates of 1930-1730 B.P. (2 Sigma

Cal.) for OZ-D and 1980-1820 (2 Sigma Cal.) for OZ-E. Though the two layers are clearly distinct stratigraphically, the dates cannot be separated statistically, indicating that the two OZs developed in a short time span, separated by a rapid accumulation of light clay deposits between them.

Occupation Zone F

Location: Trenches B, I

Extent: 34+097.5 to 34+140

Depth: 65-80 cm

Feature(s): 74, 109, 111 and 115

Estimated Date: 2770 – 2720 B.P.

OZ-F was a large, extremely dense sheet midden, appearing to be nearly paved with shell (Figure 6-25). The depth of this 12-15 cm thick zone varied from 65 to 80 cm brs due to the slope of the paleosol.

The southern part of OZ-F lies beneath OZ-D, with a clear separation of light clay between them (Figure 6-25). Within OZ-F, there were numerous areas where concentrations of charcoal and burned clay were denser than the overall concentration. However, these were poorly defined and could not be easily separated from the general high density of shell and charcoal in this zone. Thus, only four separate features (Features 74, 109, 111 and 115) were defined within OZ-F.

Occupation Zone G

Location: Trench I

Extent: 33+584.5 to 33+603

Depth: 80-110 cm

Feature(s): 84 (renamed OZ-G)

Estimated Date: 2740-2460 B.P.

OZ-G was located only in Trench I in the field, although subsequent examination of a profile of the part of Trench B dug during Phase I suggests that it may have extended south at least to 33+556. It was first identified as Feature 84 near the bottom of the trench at 110-140 cm bgs, but as its extent



Figure 6-25. The bottom of OZ F in Trench I, showing the density of shell frequently seen in this OZ.

became clear, Feature 84 was renamed OZ-G. As OZ-G continued north, however, it became increasingly shallower and thinner, until at the northern end it was only about 7-10 cm thick, 80 cm bgs, directly beneath a layer of fill and laying directly on top of Beaumont clay.

OZ-G was found throughout an unusually thick (as much as 30 cm) paleosol, with frequent scatters of shell, charcoal and burned clay. OZ-G was deepest and thickest at the south end and gradually rose and thinned as it continued north, becoming very thin at about 33+601, at which point

it was sitting directly on top of the Beaumont clay formation. During most of its extent, OZ-G was fairly dense and was much thicker than most of the Occupation Zones.

Occupation Zone H

Location: Trench L

Extent: 0 to 6 m east of centerline at 33+541

Depth: 150-155 cm

Feature(s): 125

Estimated Date: Unknown

OZ-H was located in Trench L, near the centerline of Ennis Joslin road. The depth of this zone below road surface was 150 to 155 cm, making it by far the deepest occupation zone. OZ-H was a distinct, yet relatively thin layer of broken shell, mammal bone fragments and charcoal. A single small hearth (Feature 125) was found within this zone. This zone was located at the top of a dark clay layer and the sediments that contained the occupation debris were distinctly sandier than the rest of the dark clay band.

Occupation Zone I

Location: Ennis Joslin Rd.

Extent: 33+600 to 33+608

Depth: 55-60 cm

Feature(s): 129 (renamed OZ-I)

Estimated Date: Unknown

OZ-I was observed just beneath the roadbed during scraping of the eastern half of Ennis Joslin to lower the road to the new grade, prior to addition of new road base materials. It was approximately 55–60 cm below the old road surface, and extended across most of the area being graded, although it was distinctly denser along the western edge of the area, near the centerline. It was originally designated Feature 129, but when it became clear that it covered such a large area, it was renamed OZ-I.

Most of OZ-I appeared highly disturbed, with apparently undisturbed portions only seen on the western edge, roughly

2–4 m east of the centerline. The remains of OZ-I visible just to the west of Trench B were extremely thin which explains why it was not noticed in Trench B.

Occupation Zone J

Location: Trench G

Extent: 34+012 to 34+130

Depth: 40-55 cm

Feature(s): 148-152, 154-156, 159, and 161

Estimated Date: 2480-2340 B.P.

OZ-J was located on the western side of Ennis Joslin, in the northern dune deposit. It was clearly defined within Trench G, between stations 34+012 and 34+120. When the road surface on the west side of Ennis Joslin was removed and the previously undisturbed dune surface scraped down approximately 5 cm, the surface of OZ-J was revealed across the entire western half of Ennis Joslin between approximately 34+100 and 34+130. Very large numbers of shell, scattered animal bone (most of which was fish bone), charcoal and other evidence of occupation was clearly visible in this area, and one dense concentration of shell was so well defined that it was given a feature number, Feature 161.

OZ-J and its associated features were across the street from the area in which the dense occupation debris seen in Trenches B and I were located, including OZ- B, -D, -E, and -F. In the field, it was considered possible that OZ-J is a continuation of OZ-D, based on depth below ground surface, general character of cultural materials, and the density of features in the zone. However, the radio carbon dating of samples from the five OZs made it clear that they were not related, as OZ-J was dated 2480-2340 B.P., while OZ-D was dated 1930-1730. OZ-F was dated 2770-2720 B.P.

OZ-B is directly across the street from OZ-J, but is considerably deeper below the road surface. This did not rule out a connection with OZ-J, however, because the paleosols in the northern dune deposits can be seen to descend to the east toward the False Oso mudflats. The charcoal sample from OZ-B was dated 2300-1880 B.P., making it possible, though not certain, that the two OZs are temporally related.

Occupation Zone K

Location: Trench N and Ocean Dr.

Extent: 33+558 to 33+576

Depth: 70-80 cm

Feature(s): 162

Estimated Date: Unknown

OZ-K was observed when the new free right turn from east-bound Ocean onto Ennis Joslin was graded. This is an area

that lies between the dense black clay of the pond deposits and the Beaumont Clay formation that rises from beneath the pond to the west. OZ-K is a narrow band (ca. 5 m across) that extended across the new turn lane. Although it was observed in Trench N, it was recognized at the time only as a thin scatter of shell and charcoal flecks too ephemeral to call a feature. In the context of the observations in the adjacent turn lane, however, it appears that the shell scatter seen in Trench N was part of OZ-K.

OZ-K appears to be a band of cultural materials that surrounds the northern edge of the pond. A similar deposit was not seen along the southern edge of the pond in Trench N. In fact, no evidence of cultural activity, except for an occasional shell, was seen in Trench N until OZ-K was encountered at 3+260. This may, however, be the result of previous grading of the area near the southern end of Trench N that may have removed the thin scatter of shell and charcoal seen in OZ-K. The northwestern edge of the pond deposit had been seriously disturbed by extensive previous utility construction, so that the absence of cultural deposits does not necessarily mean that they did not, at one time, exist there. In any case, OZ-K is not evidence of extensive occupation along the north edge of the pond. Although there is one area of shell in the zone that was distinct enough to define as a feature (Feature 162), the cultural material in this occupation zone was not very dense, compared to some of the occupation zones to the south on Ennis Joslin.

Summary

This chapter has provided a summary of the radiocarbon results from 41NU2. Dates show that while earlier occupations are present, the site was primarily occupied between 3000 and 900 B.P. The beginning date of this more intensive occupation is consistent with some estimates of when the barrier islands formed. An investigation of paired shell and organic dates showed that corrected, calibrated shell dates were often younger than their organic counterparts. Date ranges (2-sigma) on shell frequently did not overlap those derived from paired charcoal or organic sediment. However, a comparison of the mid-points of the two sigma ranges for the individual dates shows, with one exception, a strong relationship. We conclude that marine shell can provide accurate, precise dates in the absence of other dating materials. This may especially be the case if samples of *Crassostrea virginica* are selected for dating. We also provided basic descriptive data on the 164 features as well as 11 occupation zones documented at the site during this monitoring project. Aspects of the feature and occupation zone data, especially from dated contexts, will be used in subsequent chapters.

Chapter 7: Theoretical Framework: Modeling Hunter-Gatherer Adaptations

Steve A. Tomka

In this chapter we review the ethnohistoric record to extract some broad patterns about hunter-gatherer adaptations in the Gulf Coastal Plains and Marshes during the early sixteenth century. We look specifically at information about diet, land use, mobility and seasonal and inter- and intra-regional variability in these patterns.

Coastal Hunter-Gatherer Adaptations: The Ethnohistoric Record

Given early Spanish interest in Texas, followed by French and later Euro-American explorations into present-day Texas, a relatively large number of ethnohistoric accounts exist describing various features of the landscape, including sometimes its indigenous population. Of these accounts, the focus here will be on one source that provides details on indigenous groups specifically within the project area and the broader Gulf Coastal Plains and Marshes. We speak of Álvar Núñez Cabeza de Vaca and his account of the seven years (1528-1535) he and his companions Andrés Dorantes, Alonso del Castillo, and a Moorish slave named Estebanico spent among indigenous groups of the barrier islands and the coastal plains. Cabeza de Vaca published an account of those years, *La relacion que dio Aluar nuñez cabeça de vaca de lo acaescido en las Indias en la armada donde yua por governador Panphilo de Narbaez desde el año de veynte y siete hasta el año de treynta y seys quo bolvio a Seuilla con tres de su compañia* (usually called *la Relación*), in 1542. In 1547 a joint account by the three Spanish survivors, originally given to the Viceroy in Santo Domingo, was published in the third edition of *Historia General y Natural de los Indias* by Gonzalo Fernández de Oviedo y Valdéz (Oviedo y Valdéz 1923). We rely on the Oviedo account and the revised version of *La Relación* published in 1555. Scanned images, page by page, of the 1555 publication of *La Relación* with a new translation into English is available online (Cabeza de Vaca 1555). It is this translation that we rely on throughout this section, with page numbers reflecting the page numbers in the original 1555 Spanish edition. In addition, we also refer to T.N. and T. J Campbell's (1988) discussion of the Cabeza de Vaca account to guide us in locating the groups mentioned in the account and providing region-specific observations.

There are obvious biases in the account and, not surprisingly, detailed observations are missing on many aspects of indigenous life. In addition, it is a single account written from the perspective of a single individual with input from his surviving mates. Nonetheless, given that these individuals

lived with several indigenous groups distributed across what we assume to be the Gulf Coastal Plains and Marshes and for several years (Cabeza de Vaca 1555), the account provides us not only detail on several groups living in distinct settings but also some degree of long-term perspective that is not present in brief expedition accounts of indigenous groups.

Careful scrutiny of the *La Relación* and Oviedo's account reveals that at the time of Cabeza de Vaca's shipwreck, the Gulf Plains and Marshes region was occupied by numerous indigenous groups practicing a number of different land-use strategies. There appeared to be four land use patterns in operation during the early sixteenth century: (1) a littoral pattern with intermittent and/or regular trips to inland highly productive resource patches; (2) a river valley-based pattern with specialized trips to neighboring highly productive inland and coastal resource patches; (3) a mixed oak woodlands-prairie margins pattern with at least seasonal bison exploitation; and (4) an inland pattern devoid of perennial riverine resources. These land use patterns appear to differ in character primarily because of differences in the nature of the resources available for exploitation. We focus here on a description of the first two patterns because several historians agree that Cabeza de Vaca made landfall on or near Galveston Island and he spent the longest time with groups located either on the island or on the coast in nearby river valleys and therefore the two land use patterns have greatest relevance to our reconstructions of historic hunter-gatherer land use and mobility practices.

The Littoral Pattern

The littoral pattern appears to have been one where the bulk of the resources at least during certain seasons of the year were obtained from the beaches and waters of the Gulf of Mexico and the river valleys that empty into bays along the Gulf. Gulf resources were extracted either while on what Cabeza de Vaca (1555:39) called Malhado Island or on the shores of the mainland itself. In the case of some groups, movement into inland portions of the coastal prairies appears to have been limited or absent, at least as exemplified by Cabeza de Vaca's descriptions of the land use pattern of the groups from Malhado Island. In the case of other groups, movement into inland highly productive resource patches, particularly along river courses, might have been a common practice. For instance, Campbell (1988:74) reports that a number of coastal groups made incursions into the Central Coastal Plains region as far as Bastrop County, in pursuit of bison and probably

other terrestrial resources. Most cultural detail provided by Cabeza de Vaca refers to this land use strategy.

The shipwrecked individuals joined the island natives in early November 1528 and almost immediately experienced the effects of a strong cold front that deflated fish densities and made the harvest of plant roots difficult (Cabeza de Vaca 1555:38). The island was inhabited on occasion by native groups from at least two language groups. They each appear to have followed similar land use practices, alternating between the island and mainland camps situated along the bay (Cabeza de Vaca 1555:42). The group Cabeza de Vaca was with remained on the island from October-February. During the early part of this period they trapped fish with cane weirs and pulled roots from underwater, but later the fish were no longer abundant and the people were more heavily dependent on the roots (Cabeza de Vaca 1555:39).

At the end of February, when the roots they were eating began to sprout, the group left the island and traveled to the shores of the nearby bays using canoes (Cabeza de Vaca 1555:42). The houses built there were woven mats wrapped over bent cane poles set on piles of oyster shells (Cabeza de Vaca 1555:41). These oyster piles appear to be much like the oyster lenses or concentration noted by Martin (1930), Jackson (Jackson et al. 2004) and others found the dune sites, including 41NU2, overlooking the shore. Cabeza de Vaca (1555:41) mentions that “firewood is scarce, but mosquitoes are plentiful”. He also mentioned that fresh water was scarce during the months spent eating oysters (Cabeza de Vaca 1555:41). They stayed on the shores of the estuaries where they could procure oysters until April, then traveled back to the island to eat the large numbers of ripe blackberries available (Cabeza de Vaca 1555:41).

Early summer months were spent eating roots, and a variety of small prey including lizards, snakes, rats, and spiders. The final two months of the annual round (i.e., August-September) they live on fish and the roots of partially submerged plants.

The indigenous groups on the island are described as having animal hides for cover and using arrows that were most likely poisoned. Oviedo y Valdez (1923:23) indicates that “at times there are deer, which they kill from canoes.”

The Riverine Pattern

It appears that at least in some cases, populations living immediately inland from coastal bays were oriented towards the exploitation of inland rather than coastal resources. The Mariames and Avavares are perhaps the two best examples of this inland, riverine land use system. Cabeza de Vaca spent 18

months with the Mariames and approximately eight months among the Avavares. T.N. and T.J. Campbell (1988:25-28; and Campbell 1988:49-50) places the Mariames in the vicinity of the San Antonio and Guadalupe Rivers. It is likely that along the lower Guadalupe their territory extended to the coastal bays (i.e., San Antonio Bay) since the bays are described as being used to trap large numbers of game (Oviedo y Valdéz 1923:110).

Cabeza de Vaca’s descriptions clearly indicate that the Mariames occupied and exploited a bilobate (Campbell 1988:47) area, composed of two distinct inland segments located some 80 miles (129 km) apart and connected by a narrow travel corridor. The lower Guadalupe River portion of their range might have included portions of present-day northwest Calhoun County, northeast Refugio County, and southern Victoria County. Apparently, for nine months of the year, during the fall, winter, and spring, the Mariames inhabited and exploited the resources of this area.

Their diet consisted of a broad range of foodstuffs with some emphasis upon highly productive plant resources as they became available during the year. Pecans constituted the dominant foodstuff for one to two months in the fall season although, as Cabeza de Vaca (1555:47) mentions, pecans provide high yields only in alternate years. During years when pecan patches were highly productive, groups came from as far as 50 to 75 miles away (distances calculated based on Campbell’s [1988] reconstruction of group territories). While some of the nut meat was consumed immediately without any processing, Cabeza de Vaca also mentions that some nut meat was ground with small seeds (*granillo*) of some sort and eaten in this fashion (Cabeza de Vaca 1555:47). Although highly productive and probably capable of excess yields, no mention is made of the storage of excess nuts among the Mariames or neighboring groups.

During the winter months the primary foodstuffs consisted of the roots of two to three plants (Cabeza de Vaca 1555:52). These roots ranged from smaller than pecan to slightly larger sizes and some are described as very bitter causing the abdomen to swell. The roots were thinly distributed, hard to find, and difficult to dig (Cabeza de Vaca 1555:52). Women often foraged as far as 8 to 13 km (5 to 8 miles) from their settlements in search of these roots (Campbell 1988:50). It is not possible to discern whether these search costs are incurred within rather depleted resource patches or are due to seasonal fluctuations. In addition, to render them edible, the roots required substantial processing, including cooking for two days in some type of oven. Apparently, women invested a substantial amount of effort each night in the preparation of ovens for baking roots (Cabeza de Vaca 1555:52-53). These

high processing costs significantly reduce the post-encounter return rates for these food items and lower the ranking of such resources. Deer were hunted, both locally and by traveling out onto the prairie. The latter was done only two or three times a year because firewood and water had to be carried on these trips and mosquitoes were a serious nuisance (Cabeza de Vaca 1555:54).

Cabeza de Vaca (1555:53) claimed that intermittent food shortages were common among the Mariames, especially during the winter season. Apparently they moved every two or three days in search of food (Cabeza de Vaca 1555:53). The list of foodstuffs consumed by the Mariames includes snakes, rats and mice, various insects, some in larval forms, ant eggs, frogs, lizards, salamanders, and spiders. The bones of snakes were pulverized and eaten (Cabeza de Vaca 1555:52). It is not clear whether the diet breadth indicated by this list was a seasonal pattern reflecting starvation diets or the consumption of this range of food stuffs occurred year around.

Fish were likely obtained from the Guadalupe and San Antonio Rivers, however, it is unclear what, if any, facilities or tools were used. Apparently freshwater fish were eaten in quantity mainly during April and May following the seasonal floods when the fish could be obtained from shallow floodplain pools after the floodwaters subsided (Campbell 1988:50). Presumably the same tactic could have been employed during the fall flooding season during the second wet period of the year. Fish bones, like snake bones, were also pulverized and eaten (Campbell 1988:50).

The remainder of the year (summer) the Mariames were at the prickly pear area, placed by Campbell in the eastern portion of Duval County and Jim Wells County where the plant grew in great abundance (T.N. and T.J. Campbell 1988:9). Apparently the fruit ripened at different times for a period of about three months, between late May/early June and late August. Given that few perennial streams existed in the prickly pear grounds, the people drank ponded rainwater when it was available and prickly pear fruit juice when water was not available (Cabeza de Vaca 1555:56). The prickly pear fruit was eaten immediately after it was picked or shortly thereafter. Unlike in the case of the pecans, however, there is mention of some of the prickly-pear fruit being dried for consumption during the return trip to the lower Guadalupe River. The juice of the fruit was squeezed out, the fruit was opened, dried in the sun, and thereafter carried in some kind of flexible container. Mention is also made of the peelings of the prickly pear being ground into a powder (Cabeza de Vaca 1555:53). Although much of the discussion about prickly pear use focuses on this fruit, Oviedo also mentions that

snails were eaten while the Mariames were in the prickly pear grounds “snails were searched for carefully” (Oviedo y Valdez 1923).

Although there is no mention of deer being hunted while in the prickly pear area, it is likely that if the large congregation of people did not frighten away the deer those that were encountered were hunted while in the region. What is clear is that immediately preceding the move to the prickly pear area, and before moving away from the coastal bays the Mariames organized a communal deer drive during which large number of deer and other animals trapped in front of the drivers were herded into the bay and kept there until they drowned and washed ashore. This method of hunting tended to yield large number of animals. In one case 60 men took between 200-300 deer, while at another occasion a total of 500 deer were taken (T.N. and T.J. Campbell 1988:18). On the other hand, some communal hunts resulted in few or no deer. It is possible that if this hunting technique was timed to take place immediately before the move to the prickly pear area, it was designed to provision the population with meat while in this patch for 1.5 to 3 months.

The Avavares are the third group with which Cabeza de Vaca spends extended time. They are briefly mentioned here because they may be the group that was most closely associated with the project area. According to Campbell, the Avavares may have inhabited the area between present-day Lake Corpus Christi and Copano Bay and possibly adjacent areas in Bee and Refugio counties northwest of Corpus Christi Bay. Other researchers place the Avavares on both sides of the lower Nueces River in parts of Duval, Jim Wells, and San Patricio Counties. They were an inland group living west and immediately inland from the coastal Quitoles (T.N. and T.J. Campbell 1988:25). They appear to represent a variant of the riverine pattern with perhaps lower reliance on coastal resources compared to the Mariames.

Based on Cabeza de Vaca’s and Oviedo’s descriptions, it is clear that at least during the summer the Avavares lived in or immediately adjacent to the prickly pear area. One cannot make a clear distinction between their summer and winter ranges indicating perhaps that they inhabited the same general area throughout the year. Other than prickly pear, the record is unclear what resources were exploited by these people. There is specific mention that they did not have fish, or collect pecans or acorns (T.N. and T.J. Campbell 1988:26). However, since one of the seasons is identified as the time when the fish die, it is puzzling why fish are not mentioned as one of the available resources. This would suggest that they did inhabit a river valley (e.g., the lower Nueces River), although if this was the case, it would be puzzling that no other

river bottom resources are mentioned as being systematically exploited (T.N. and T.J Campbell 1988:26).

During six of the eight months (Sept. 1534-June 1535) Cabeza de Vaca and his fellow Spaniards were with the Avavares they all experienced severe shortages of food and suffered hunger. Winters were apparently particularly difficult in terms of foodstuff availability. Much of the summer subsistence was geared towards the exploitation of prickly pear. In late summer they are described as being in the prickly pear patch searching for prickly pear that escaped harvest by other groups. On another occasion when several days of searching produced no ripe fruit, the Avavares went to a stream valley, to gather seed pods from trees whose pods the Spanish described as like “*hieros*” (vetches or lentils). This tree has been identified as the Texas ebony (*Ebenopsis ebano*), which holds its seedpods into the winter season (T.N. and T.J. Campbell 1988:26). According to Oviedo, during the winter months, the Avavares ate mainly roots. (Oviedo y Valdéz 1923). Deer were hunted and Cabeza de Vaca mentions that at one point the Avavares gave them so much venison that “we did not know where to put the meat” (Cabeza de Vaca 1555:58). It is not known whether the meat was sun dried or other wise preserved or fresh. It also is unclear as to whether the food scarcity described among the Avavares is due to the likelihood that they inhabited more xeric settings or whether it was the product of the typical seasonal fluctuation in resources.

Apparently, their annual land use system was closely tied to the succession of the seasons and the various foodstuffs that became available. It is perhaps because the lands around the Avavares are described as having few resources, that mobility was relatively high among them. Their houses are said to have been like those of the Mariames, a bent-pole framework covered by mats. When they moved their camps, as often as every three days, they dismantled it and transported it to the new location (Cabeza de Vaca 1555; Oviedo y Valdéz 1923). Again, it would be great to know if the frequency in moves is the typical year round pattern or applies only to a particular season of the year (i.e., very early or late summer).

The previous review of historic period adaptations by indigenous groups along the Gulf coast provides a surprising amount of information on the subsistence and mobility of hunter-gatherers living within or in the vicinity of the project area. We have learned that the Middle Gulf Coast was inhabited by a large number of groups positioned adjacent to each other along the coastal belt and immediately inland. Some portions of the landscape and resources appear to have been shared by multiple groups (i.e., prickly pears and pecans) while access to others may have been restricted to specific groups (i.e.,

coastal strips). Mobility was in general high and included both logistical (i.e., deer and/or bison hunting parties on the prairie) and residential moves (i.e., camp movements every three days). Daily foraging rounds may have varied between 8-13 km and long distance travel to shared resource patches may have encompassed as much as 113-129 km (figures calculated based on group locations identified by T.N. and T.J Campbell (1988) based on the Cabeza de Vaca account [version unknown]). Hunting may have focused on medium body size game acquired either through communal efforts, small logistically organized hunting parties or individual pursuit hunting. Some long distance hunting expeditions may have been undertaken to pursue bison. Campbell (1988) for instance, surmises that some coastal groups may have penetrated far inland to hunt bison on a seasonal basis. Plant resource acquired included roots and tubers, prickly pear, pecans, and seeds. Oysters, fish, waterfowl, and a variety of amphibians, reptiles and rodents supplemented the diet. While this is a welcome amount of detail, that serves as an important window into the past, as with any historical account, it does not help us understand why land use and subsistence practices were organized the way they appeared to have been. To begin to define what were the broad conditioning factors underlying this ethnohistoric landscape and how they operated, we turn to models provided by human behavioral ecology.

Models from Behavioral Ecology

The theoretical position we adopt in this document to investigate prehistoric hunter-gather adaptations on the Gulf Coastal Plains and Marshes is human behavioral ecology. Human behavioral ecology (HBE) is a subset of evolutionary ecology that focuses on the explanation of variability in human behavior through the principles of natural selection. Its starting position is that human behavior is the product of a number of evolutionary mechanisms and processes operating at different temporal and spatial scales. That is, while a wide range of behaviors exists within the human repertoire at any one time, over time, natural selection will favor those sets of behaviors that tend to be optimal or more precisely optimal within certain constraints (Foley 1985; Winterhalder and Kennett 2006:11). The constraints derive from the fact that a multitude of factors such as lack of full information, lag time between resource states and behavior adjustments, idiosyncratic differences in skill levels and cognitive differences between individuals, result in less than perfect match or synchronicity between resource states and hunter-gatherer behavior.

Four fundamental concepts underlie the optimization premise: (1) any given behavioral alternative has opportunity costs; (2) all human behavior has some degree of risk; (3) continuous

and/or repeating human behavior has a marginal value; and (4) rewards for future behaviors with uncertain outcomes have lesser value than behaviors with immediate and certain outcomes (Winterhalder and Kennen 2006). Employing these underlying premises, the intention of HBE analysis is to determine through cost/benefit analysis the best combination of alternative sets of resource procurement behaviors given a set of constraints (e.g., resources available at a given time, technological means, etc), an assumed set of goals (e.g., maximize energy capture, minimize risk), as measured using a given currency (e.g., Kcal, time, mating success, prestige, etc). Because optimization modeling is a key framework used by HBE studies to develop expectations regarding hunter-gatherer behavior, we briefly review some optimal foraging models and propose that one in particular, the habitat choice model, with some modifications is appropriate for the study of prehistoric hunter-gatherer foraging strategies in the Gulf Coastal Plains and Marshes.

General Background to Optimal Foraging Models

In Chapter 2, we identified the Gulf coastal strip as a highly diverse mosaic of resources. Within this mosaic, we have defined five relatively homogenous habitats each with its own mix of biological communities (see Chapter 2). Seasonal changes and long-term modifications in climatic conditions, coupled with unpredictable weather episodes strongly influence the types of resources available within each habitat, their seasonal or long-term availability, their relative abundance, and their nutritional condition. Specific aspects of the resources present in these habitats and fluctuations in their availability are discussed in greater detail in the next chapter.

Among hunter-gatherers, problems created by spatial variation in resources are commonly solved by changes in mobility strategies and adjustments in group size (e.g., fissioning and aggregating; Kelly 1995:213-216). Mobility strategies have several components that can vary, including the frequency of moves, the distance moved, and the organization (e.g., logistically organized task groups, higher residential mobility) of the moves. Similarly, the temporal fluctuations in resources, including seasonal changes in resource availability and quality (e.g., molluscs, seeds, nuts), year-to-year fluctuations in resource density (i.e., nuts), and longer-term changes in resource structure (e.g., increases or shrinkages in arboreal vegetation communities vis a vis grasslands), are commonly solved by changes in diet, shifts in mobility strategies, fluctuations in group composition and size, and technological adjustments including the use of storage.

Since the 1970s, optimal foraging models adapted by anthropologists and archeologists from animal behavioral

ecology have been one of the main approaches used to model hunter-gatherer behavior. Since the tendency for optimal behavior is conditioned by long-term evolutionary forces, optimal foraging theory has become the principal methodological tool employed by human behavioral ecologists to study prehistoric hunter-gatherer adaptations. Two principal models, a "Prey Choice Model" which was explicitly developed for a single predator pursuing a single prey in a homogenous environment (Charnov et al. 1976; MacArthur and Pianka 1966; Winterhalder 1981), and a "Patch Choice Model," designed for situations where multiple prey are clustered in a more heterogeneous setting (Charnov 1976; Kaplan and Hill 1992), have been developed to account for foraging behaviors.

As we have argued elsewhere (see Tomka et al. 2004), while many assumptions of these foraging models are clearly violated by hunter-gatherers, and while parameters specified for analysis (e.g., caloric content; post-encounter return rate) are often difficult to estimate in archeological situations, we find the models appealing as they provide an explicit cost/benefit framework for analysis. Also, we recognize that by their very nature these models are simplified versions of more complex hunter-gatherer behaviors. Models are simplified versions of reality with understood and often quantified relationships between conditioning factors and behavioral outcomes. We compare the material signatures and adaptive stances reconstructed from the archeological record to these models to search for agreements and disagreements between them. These comparisons often result in refinements of the models and greater understanding of prehistoric hunter-gatherer foraging patterns.

Applications of foraging models (see Stephens and Krebs 1986) in hunter-gatherer studies typically quantify *costs* as time expended in searching for, pursuing, capturing, and processing food items and the *benefits* as energy (kilocalories [kcal]) obtained from the acquired food (but see Jochim 1981; Speth and Spielmann 1983). Maximization models assume that foragers attempt to maximize average return rates in the context of different cost/benefit outcomes for different prey. Search costs reflect the amount of time spent looking for game or resource patches, and handling costs include the amount of time required to pursue, capture, and process foods. For hunters and gatherers, we also include in handling costs time associated with production of tools (e.g., raw material acquisition, tool shaping, maintenance, production of ceramics; Bousman 1993), and the manufacture of facilities and features employed in resource processing (e.g., hearths, hot rock ovens).

A critical element of all foraging models is the ranking of prey alternatives. Potential prey items are ranked in terms of

post-encounter return rates (i.e., Kcal/hr). For animals, this ranking is strongly correlated with body size since larger-bodied animals (e.g., mammoths, bison) are more profitable (higher returns relative to handling costs) than smaller-sized animals such as a deer, rabbit or mouse (see Kaplan and Hill 1992: Figure 6.1; Simms 1987: Table 5). Even though some plant resources have higher caloric content than meat, plants usually rank below animals primarily because of their smaller package size, though there are exceptions. This general trend in ranking can be seen in Figure 7-1. Here we present box plots for a series of post-encounter return rates gathered from a variety of experimental and ethnographic sources (see Kelly 1995; Simms 1987). The rates have been grouped by body size for animals and roughly by seed size for plants. Clearly, when seen strictly in terms of handling costs, the return rates on large mammals in this example are extremely high, while the return rates for collecting and processing small seeds are extremely low.

Although, as noted, search costs, are not taken into account in prey profitability rankings, these costs do play a critical role in determining the actual diet of hunter-gatherers since they represent energy expended in capturing a prey and therefore

enter into the calculation of the overall energy balance of the individual. Because search costs are treated separately from handling costs, it is possible to study the effect of prey abundance or changes in search techniques on prey selection. In the classic Prey Choice Model, as more resource types are added to the diet, search costs decline because resources are more frequently encountered. However, these new resources, being lower ranked, have higher handling costs and/or lower caloric benefits (lower profitability). Foraging models predict a tradeoff, then, between handling cost, energy benefits, and search costs that will maximize the average return and produce an optimal diet (see Figure 7-2). These models predict that foragers will continue to add lower-ranked resources to the diet, increasing the diet breadth, so long as the overall profitability of the diet, seen in terms of total costs and benefits, is increased. Furthermore, resource types should be dropped from the diet, reducing the diet breadth, when doing so would increase overall profitability.

From this perspective, the decision to include or exclude a given resource type is determined not by the density or profitability (i.e., net energy gain per pursuit/handling cost per pursuit) of that resource type, but by the density and profitability of

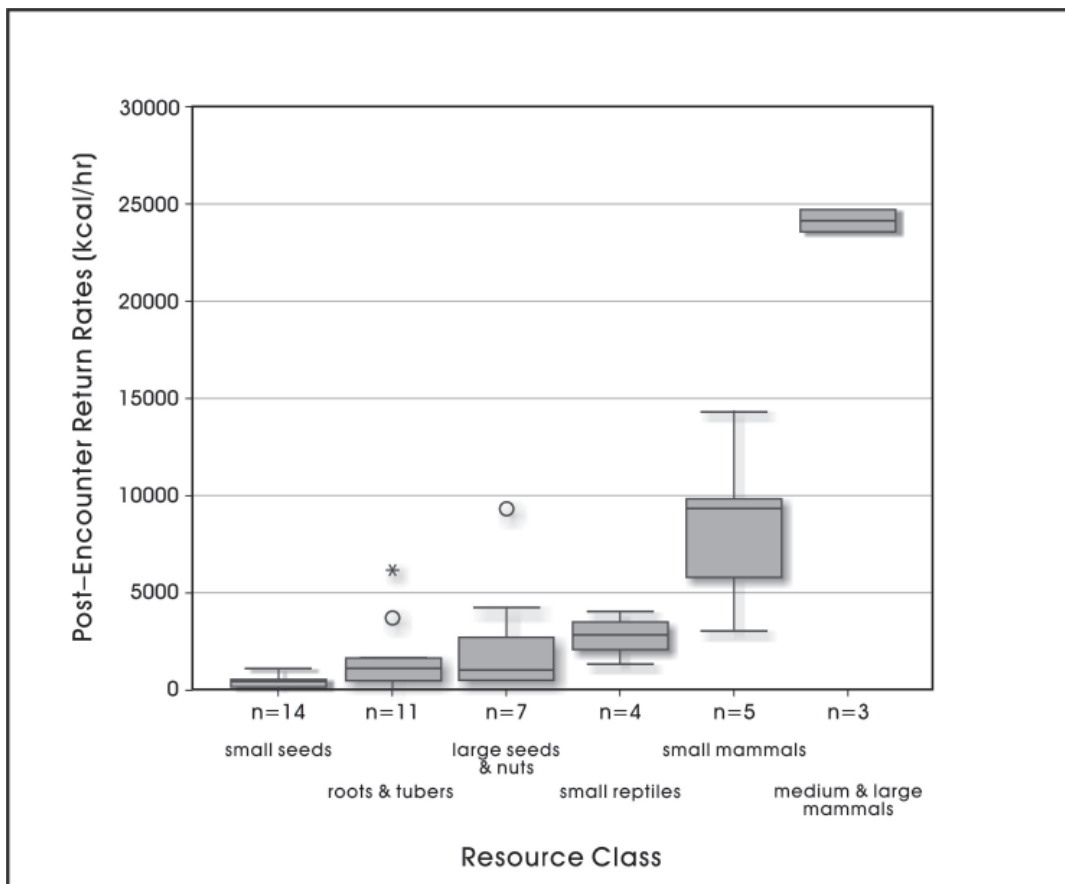


Figure 7-1. Return rates by resource class (from Cane 1987; Kelly 1995; Simms 1987).

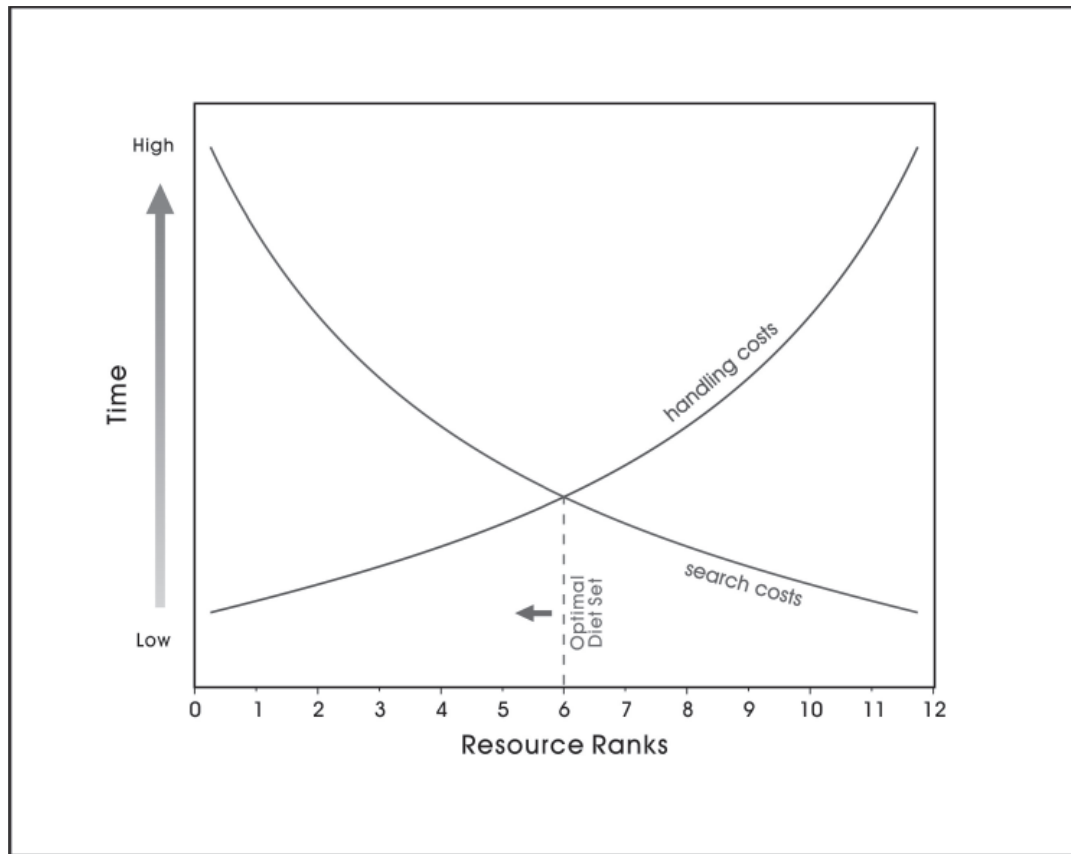


Figure 7-2. Optimal diet set as defined in Prey Model (after MacArthur and Pianka 1966).

higher ranked prey items. In situations where the frequency of higher ranked items increases, lower ranked resources could be dropped from the diet, thus increasing the average return.

While the Prey Choice Model assumes that the goal of the forager is to maximize the rate of energy intake, a substantial body of cross-cultural data suggests that human male and female foragers target different resources on their daily foraging trips (see Bird 1999). Female foragers tend to harvest prey packages that are smaller than those pursued by men, have a relatively low risk of pursuit failure, and often consist of resources that have high processing costs (Bird 1999:66). In contrast, men pursue prey that are larger, have a higher risk of failure, and are associated with lower processing costs (Bird 1999:66). The two foraging tactics are partially overlapping and highly complementary in having the overall effect of dampening some day-to-day fluctuations that may otherwise result from the high failure rates of hunting large- or medium-size prey. This contrast between model expectations and behavior suggests that actual hunter-gatherer behavior may be conditioned by a combination of factors (i.e., rate maximization, risk minimization, and others) particularly in environmental contexts where the daily caloric intake may be subject to significant variability (Levins 1968:10-38).

Over time and as a product of repeated applications of optimization approaches in both animal and human ecology, new models have been developed to address a broad range of resource procurement behaviors and models have become more capable of addressing real-world complex scenarios. For instance, the Patch Choice Model relies on the marginal value theorem to highlight that harvests will continually decline with increasing patch residence time leading to an inverse relationship between patch residence time and return rates over time (Stephens and Krebs 1986). The model predicts that a given forager will leave a patch when the rate of return in that patch has dropped to the mean return rate for all other patches that are available for exploitation. Two possible patch utilization approaches may be optimal under this model. The first consists of short patch residence time targeting the harvest of the most profitable prey accompanied by rapid movement between patches as the most profitable resources are depleted in each patch. The second approach involves longer residence in each patch while resources are utilized to the point of near sub-optimality followed by longer travel distances to new patches because a larger area surrounding the residential camp has been depleted of resources.

The Central Place Foraging Model also deals with forager decision making relative to multiple competing resource

patches exploited from a centrally located camp (Orians and Pearson 1979). Rather than narrowly limiting foraging behavior to sequential foraging from one patch to another, this model allowed for the exploitation of multiple patches from the same central location. The critical variable introduced in this model is the role of roundtrip travel and transportation costs. The expectation derived from this model is that as roundtrip travel costs increase, in particular as a result of successfully acquired prey or resources, the forager should become more and more selective about what he/she will harvest. A number of studies have used this model to explain the differential transport of resources from hunting or procurement localities to the base camp (Bettinger and Malhi 1997; Bird and Bliege Bird 1997; Metcalfe and Barlow 1992). Others (Brannan 1992; Jones and Madsen 1989; Rhode 1990) have attempted to model the effect of resource transportation costs on settlement location vis a vis resource patches with distinct post-encounter return rates.

For instance, Zeanah (2000) considers the issues of where to locate a camp when faced with the choice of two highly ranked resources (i.e., large mammal versus hyper-productive seed patches) located at some distance from each other yet harvested during the same season. The quantitative model suggests that transportation costs have a significant impact on the optimal solution for camp location. Specifically, when the contribution of a highly ranked target resource to the diet of a forager is high, and the prey is profitable due to frequent encounter rates, the forager should reside in close proximity to that target prey. Close proximity and high encounter rates should ensure high after-transport profitability. However, as encounter rates decrease, the after-transport profitability of the target prey also will begin the decrease ultimately leading to a widening of the diet breath to include the next highest ranked resource. If the contribution of this new resource to the caloric intake of the forager increases beyond that of the first resource due to higher after-transport profitability, the forager should relocate in its proximity. That is, the relative ranking of the two camp locations should switch Zeanah (2000:11). Of the two, the prey that makes up the lower contribution to the diet may still be procured through logistical organization while daily foraging trips from a central place camp should be used to access the prey that represents the highest contribution to the diet.

Oneway in which hunter-gatherers may decrease transportation costs is through field processing. However, field processing itself has a caloric cost that can decrease the net caloric gain from a target prey when included in prey profitability calculations. In the end, while field processing may reduce transportation costs through lighter loads, it may severely lower the transportation threshold for given prey species due to its energetic costs. In considering the field processing of acorns and mussel shell, Bettinger and Malhi (1997)

conclude that high field-processing costs severely deflate the transport threshold and for acorns only minimal field processing should be carried out otherwise a negative energy balance will quickly be reached. In contrast, the procurement and transportation of some mussels should involve no field processing to maintain a positive energy balance while the exploitation of other species may be profitable even following complete field processing. The choice between these two alternatives is conditioned by roundtrip travel distance, the productivity of different mussel colonies (i.e., abundance of mussels and size of meat-package), the caloric yield from different mussel species and the costs of field processing the different species. These same factors underlie the differences in optimal field processing-transportation strategies between acorns and mussels.

Responses to changing diet breadth can elicit changes not only in the frequencies and distances one moves between resource patches but also the overall organization of mobility. Binford (1980) suggested that new insights may be gained regarding land use and technological organization by viewing hunter-gatherer mobility systems as representing an organizational continuum encompassing “foragers” and “collectors”. Collector resource procurement strategies have low residential mobility and rely on task-specific groups sent from central camps to obtain distant resources in settings with high temporal or spatial disparity in resources. Foragers, frequently move residential locations to acquire food within environments characterized by ubiquitous, low-density resources (Binford 1980; Kelly 1995). Bettinger and Malhi (1997:897) note that field processing and transportation costs can significantly influence these organizational strategies. Specifically, they suggest that logistically organized procurement that involves field processing is in general a costlier strategy than the procurement of the same resource through residential mobility where foragers map onto a productive resource. They also suggest that as residential mobility decreases and hunter-gatherers become increasingly more tethered to staple resources, logistical procurement strategies will be employed to take lower ranked resources less profitably (Bettinger and Malhi 1997:897).

The final optimization model mentioned here and discussed in more detail in the next section is the Habitat Selection Model (otherwise known as the ideal free distribution model; Kennett et al. 2006; Sutherland 1996). The underlying assumption of this model is that individual foragers and foraging populations should choose to occupy the patch where suitability (i.e., fitness, breeding success, or rate of intake) is highest (Sutherland 1996:14). In the model, patches are defined as more or less heterogeneous aggregates of resources and the forager is faced with deciding where to locate. The quality of the habitat depends on resource abundance and the

density of the population inhabiting and using it (Winterhalder and Kennett 2006:16). The model predicts that habitats will be occupied in their rank order and also factors in that certain human behaviors may have advantageous consequences (e.g., creation of clearings) while other behaviors may have negative consequences on the ranking of the habitat (e.g., settlement at high population densities and resulting resource depletion).

Several key assumptions underlie the ideal free distribution model (Sutherland 1996). The first of these assumptions is that as the number of competitors increases within a patch, the suitability (read return rate) of that patch declines (Sutherland 1996:9) as a result of negative feedback. Interference is one form of negative feedback and results in short-term decline in intake rate due to the presence of others in the patch who may interfere with a forager's ability to carry out optimal solutions (Sutherland 1996:15-19; i.e., conflicts among neighboring groups may prevent access to desired patches). Population dynamics suggest that with weak interference foragers will tend to aggregate within the habitats with highest prey densities, however, as forager population densities increase, increased interference will lead to a more even distribution of the foraging population into other habitats resulting in a relatively even intake rate in all habitats (Sutherland 1996:17). Depletion, the second form of negative feedback, refers to the actual removal of prey, and can clearly impact potential return rates within a habitat (Sutherland 1996:32-37). Modeling of depletion dynamics suggests that while foragers will always aggregate into the most productive habitats (i.e., the one with the highest density of prey), as prey densities are depleted to levels equaling those of the next richest habitat, foragers will begin to use both habitats. The ideal free distribution also assumes that patches may differ in quality (i.e., suitability, return rate) and the manner in which they are used depends upon the density of competitors (Sutherland 1996:7). In species with territoriality or dominance hierarchies, the ideal distribution no longer applies because individuals are not free to move between patches. However, the concept of the ideal despotic distribution may be used to describe the expected distribution between patches (Sutherland 1996:9-10). The ideal despotic distribution recognizes that the first forager to occupy a habitat may gain higher return rates than later arrivals. Therefore, the suitability (read return rate) of the patch or habitat declines for each new forager arriving into the patch. However, the newly arriving foragers do not necessarily decrease the suitability of the patch for those foragers that were already there because through dominance or territorial behavior the original settlers can or may maintain exclusive control of their patch.

The model has been used to investigate population dynamics and animal behavior (Sutherland 1996), and trends in the colonization of Oceania (Kennett et al. 2006). One

particularly useful aspect of the model is that while most applications of Optimal Foraging Theory have been at the scale of the individual prey or patch, the habitat selection model can be applied with equal interpretive success to larger analytical units such as a habitat and allows the study of not only individual but population foraging behavior. We propose to look at hunter-gatherer adaptations at this scale because we feel that it more accurately reflects the scale at which decisions regarding subsistence and mobility are managed among hunter-gatherers.

In addition, the model expressly takes into account not only prey ranking but also the role of certain population dynamics such as interference, depletion, and the buffer effect to generate expectations regarding patch use and residence time. The inclusion of population dynamics in turn broadens the range of phenomena that can be modeled or studied well beyond subsistence and settlement practices and allows forces other than climatic phenomena to impart changes in resource density and distribution, and patch rank. It is because of these features and its close theoretical relationship to the broad family of optimal foraging models that we propose to employ it to model coastal hunter-gatherer land-use and subsistence strategies, particularly as they are manifest at the 41NU2 site.

Technological Adjustments

In addition to changing diet breadth in an effort to stay within the optimal zone, technological responses may also be taken to maximize the average return rate. For example, the abandonment of complex, expensive technologies in favor of simpler, less costly processing methods, or abandonment of specialized search methods, may be a viable option under some conditions of increasing frequency of higher ranked resources. Conversely, decreasing abundance of highly ranked prey may lead to a series of technological responses such as: (1) the adoption of more expensive technology in order to increase the kill ratio, and thereby increase the profitability of the less frequently encountered higher ranked prey; (2) the use of technologies that will allow more intensive processing of the prey thereby extracting more nutrients than was previously possible; (3) the shift in technology that increases the encounter rate of that prey type (i.e., snow mobiles). Search strategies may also shift to alternative resources with similar return rates.

Technological adjustments that can aid in increasing return rates to make a resource more profitable can also be made by major shifts in the organization of procurement. For instance, game drives can result in the aggregation of large numbers of highly ranked prey resulting in high return rates even when game population densities are low. The

ethnohistoric mentions of fire-aided game drives on the Gulf Coast may actually be examples of this type of technological adjustment to low game densities. Similarly, if the density of lower ranked prey increases (anchovies and other schooling fish) certain capture methods and facilities can result in the acquisition of large numbers of small meat packages that when treated as a single unit captured during a single even will significantly raise the return rate for that food resource (Madsen and Schmitt 1998). Facilities and tactics such as nets, weirs, and fish poisoning can serve to increase capture rates for a given lower ranked food resource. The increase in return rates with increasing abundance of otherwise low-ranked resources derives from dramatic reductions in pursuit time. The en-mass capture of large numbers of small prey items (e.g., a basket full of anchovies) with not much more than the same effort that it would take to capture a single prey item, results in a significant increase in calories contributed without a proportional increase in calories expended in securing the prey items therefore resulting in a significant increase in post-encounter return rates. As Ugan (2005) and Ugan et al. 2003, show, however, such gains would be realized only if the procurement technology is not overly expensive to manufacture and there is minimal to no energy expended in processing the harvest.

Summary

In this chapter, we have reviewed data on sixteenth century hunter-gatherers as they were described in the written records produced by Cabeza de Vaca (1555) and Oviedo y Valdéz (1923), one of his companions. We also relied on Campbell (1988) and his research related to the reconstruction of the route taken by Cabeza de Vaca on his way back to Pánuco. We realize that much of the picture presented above relies

on best-guess reconstructions of where the shipwrecked party landed, where it came to shore and the routes of their travels. Similarly, the locations of the named groups and the regional mobility patterns are best-guess estimates. Nonetheless, few historians would argue that the hunter-gatherers described lived somewhere on the coastal stretch and utilized both marine, freshwater and terrestrial resources. Therefore, the picture that we are able to extract from the accounts mentioned is an example of hunter-gatherer adaptations to resource structures that are not unlike what we would expect Gulf Coastal groups would have faced during the Late Prehistoric and Historic Periods, if not even earlier.

The ethnohistoric data suggests a high population density packed along the coastal corridor. Several resource extraction strategies appear to have been practiced by these groups. We focused on the detailed discussion of two of them, the littoral and riverine pattern. Following the review of the ethnohistoric record, we looked at optimal foraging models derived from human behavioral ecology to begin assembling the theoretical toolkit that will allow us to generate expectation regarding prehistoric hunter-gatherer land use and subsistence.

We outlined the underlying assumptions of the prey selection, patch choice, central place foraging models and proposed that aspects of all of them can be used in conjunction with the habitat selection model to construct expectations regarding the positioning of hunter-gatherer populations across the landscape, their mobility and diet breadth, as well as prehistoric population dynamics within the Corpus Christi area. In the next chapter (Chapter 8) we summarize the resource base and its spatial and temporal structure to provide a foundation that will anchor specific expectations regarding hunter-gatherer adaptations that we propose in Chapter 9.

Chapter 8: Resource Structure and Nutrition

Steve A. Tomka and Jennifer A. Thompson

In this chapter we describe the broad range of resources that would likely have been available to prehistoric hunter-gatherers of the coast and coastal plains and provide nutritional information from published sources for some species. In addition to compiling an inventory of potential resources, our initial goal was to also pull together information on nutrition and population dynamics for the selected resources. This search was intended as a pilot project to determine what was and what was not available in terms of data types that would allow archeologists to pursue quantified research in optimal foraging and human behavioral ecology.

The coast represents a rich mosaic of niches with unique but complementary resources for plant and animal species. Since hunter-gatherer groups would have used both freshwater and marine aquatic and terrestrial resources, our review will include the range of potentially available resources within both environments. Our goal in providing this review is to identify the principal potential resources and define their spatial and temporal structure and the factors that may cause predictable or unpredictable fluctuations in their availability. However, knowledge of variations in available amounts of resources will not necessarily provide an explanation of subsistence variation or relative dietary importance of any single species. Many have turned to economic analysis of resource procurement calculating net energy and time costs for hunting and processing in order to explore unexpected relationships between resource abundance, the assumed dietary importance, and the archeological record (Reidhead 1976, Perlman 1976, Osborn 1977, Glassow and Wilcoxon

1988). Any such economic study should include nutritive values per unit of weight as well as fluctuation between age, sex, and season of available resources. However, this data is difficult to come by. Many case studies refer to Watt and Merrill (1963) for nutrition information, finding the nearest applicable item listed (Glassow and Wilcoxon 1988, Reidhead 1981), submit samples for nutritional analysis as part of the research, or rely on other published sources and adjusting data to include factors that would effect the nutritional gain of a given resource. Relying on general data such as from Watt and Merrill (1963) disregards specific nutrition information at the species level, age group, season, and habitat of the resource, a sacrifice most researchers find prudent as finding sources reporting nutrition information specific to seasonal or age fluctuations is even more problematic. Submitting samples of possible resources for nutrition analysis was not possible for this study. Therefore, the literature search serves to assess the availability nutritional data at the species level for resources available prehistorically along the Central Texas Coast.

Terrestrial Mammals, Reptiles, and Amphibians

Table 8-1 presents the 51 mammalian species found within the Coastal Prairies sub-region of the Gulf Coastal Plains, including the areas surrounding the numerous bays of the Gulf. The table also includes a column describing the preferred settings within which the species is encountered. The data were compiled from the mammalian species distribution

Table 8-1. Mammalian Species Found in Terrestrial Settings of the Gulf Coastal Plains and Marshes

Size	Taxa/Species	Weight (g)	Found In	Calories (or Kcals)	Protein Content (kg)	Unit Size	Reference
Large	Bison	600000	prairies/open fields; open woodlands?;	–	–	–	–
	Mountain Lion	157250	sandstone/limestone ledges, bluffs; thickets; wooded slopes; open woodlands;	–	–	–	–
	Black Bear	125000	open woodlands; wooded slopes; glades;	134.29	23	100 kg units	Reidhead 1981
Medium	Jaguar	52800	open woodlands; thickets;	214.24	20.6	100 kg units	Reidhead 1981
	White-tailed Deer	50000	open woodlands; glades; thickets;disturbed grounds;	334.17	17	100 kg units	Reidhead 1981
	Pronghorn	45000	prairies/open fields; xeric habitats;	254.22	19.4	100 kg units	Reidhead 1981
	Red Wolf	30000	open woodlands;wooded slopes; prairies/open fields; glades;	–	–	–	–
	Collared Peccary	19000	thickets; xeric habitats; barrens;	–	–	–	–
	American Beaver	18000	marshes, sloughs, ponds; alluvial stream bottoms;	125.75	21	100 kg units	Reidhead 1981
	Coyote	17000	open woodlands; glades; thickets; wooded slopes; prairies/open fields; disturbed grounds; xeric habitats;	138.43	20.8	100 kg units	Reidhead 1981

Table 8-1. Continued...

Size	Taxa/Species	Weight (g)	Found In	Calories (or Kcals)	Protein Content (kg)	Unit Size	Reference
Small	Common Raccoon	8500	open woodlands; glades; wooded slopes; marshes, slough, ponds;alluvial stream bottoms;	255.55	17.6	100 kg units	Reidhead 1981
	Porcupine	8000	sandstone/limestone ledges, bluffs; wooded slopes; open woodlands;	198.34	19.2	100 kg units	Reidhead 1981
	American Badger	7000	prairies/open fields; xeric habitats - occupy entire range for prairie dogs and ground squirrels;	1258	–	1 kg	Simms 1984
	Bobcat	7000	sandstone/limestone ledges, bluffs; thickets; wooded slopes; open woodlands;	–	–	–	–
	River Otter	6500	marshes, sloughs, ponds; alluvial stream bottoms;	–	–	–	–
	White-Nosed Coati	4500	open woodlands; wooded slopes; sandstone/limestone ledges, bluffs;	200.09	21.9	100 kg units	Reidhead 1981
	Common Gray Fox	4000	open woodlands; thickets; wooded slopes; xeric habitats;	–	–	–	–
	Stripped Skunk	4000	open woodlands; glades; wooded slopes; thickets;	–	–	–	–
	Black-tailed Jack Rabbit	3250	xeric habitats; barrens; sandstone/limestone ledges, bluffs;	134.29	23	100 kg units	Reidhead 1981
	Opossum	3150	open woodlands; glades; prairies/open fields; marshes, sloughs, ponds;alluvial stream bottoms;	373.67	15.9	100 kg units	Reidhead 1981
	Swamp Rabbits	2250	marshes, sloughs, ponds; alluvial stream bottoms;	441.37	13.8	100 kg units	Reidhead 1981
	Common Hog-nosed Skunk	1900	open woodlands; glades; wooded slopes; thickets;	326.43	17.3	100 kg units	Reidhead 1981
	Eastern Cottontail	1500	open woodlands; disturbed grounds; glades; open fields;	–	–	–	–
	Black-tailed Prairie Dog	1500	prairies/open fields;	–	–	–	–
	Ringtail	1250	sandstone/limestone ledges, bluff; wooded slope/slope bases; barrens;	–	–	–	–
	Eastern Fox Squirrel	950	open woodlands; glades; alluvial stream bottoms; wooded slopes/slope bases;	–	–	–	–
	Common Muskrat	881	marshes, sloughs, ponds; alluvial stream bottoms;	–	–	–	–
	Mink	782	alluvial stream bottoms; marshes, sloughs, ponds;	–	–	–	–
	Eastern Spotted Skunk	565	open woodlands; glades; prairies/open fields; wooded slopes; sandstone/limestone ledges, bluffs;	–	–	–	–
Eastern Gray Squirrel	445.6	open woodlands; thickets; marshes?;	1078	–	1 kg	Simms 1984	
Long-tailed Weasel	400	coextensive with pocket gophers and ground squirrels;	–	–	–	–	
Very Small	Eastern Woodrat	275	alluvial stream bottoms; marshes, sloughs,ponds;open woodlands; sandstone/limestone ledges, bluffs, prairies/open fields;	134.29	23	100 kg units	Reidhead 1981
	Southern Plains Woodrat	257.5	thickets; xeric habitats; barrens;	169.71	22	100 kg units	Reidhead 1981
	Mexican Ground Squirrel	213	prairies/open fields; xeric habitats;	262.38	19.2	100 kg units	Reidhead 1981
	Thirteen-lined Ground Squirrel	157	prairies/open fields;	207.7	20.9	100 kg units	Reidhead 1981
	Hispid Cotton Rat	115	prairies/open fields; barrens; disturbed grounds;	–	–	–	–
	Eastern Mole	75	prairies/open fields; glades; open woodlands;	–	–	–	–
	Eastern Flying Squirrel	54	open woodlands; wooded slopes/slope bases;	135	–	100 g raw	Watt and Merrill 1963
	Gulf Coast Kangaroo Rat	52	barrens; disturbed grounds;	1078	–	1 kg	Simms 1984

Table 8-1. Continued...

Size	Taxa/Species	Weight (g)	Found In	Calories (or Kcals)	Protein Content (kg)	Unit Size	Reference
Very Small	Marsh Rice Rat	51	marshes, sloughs, ponds; open fields; rarely open woodlands;	-	-	-	-
	Cotton Mouse	42.5	open woodlands; glades;	-	-	-	-
	Hispid Pocket Mouse	38.5	open woodlands; glades; wooded slopes; disturbed grounds;	-	-	-	-
	Southern Short-tailed Shrew	23	open woodlands; glades; prairies/open fields;	117.44	19.9	100 kg units	Reidhead 1981
	Deer Mouse	22.5	prairies/open fields; open woodlands; xeric habitats; disturbed grounds; wooded slopes;	-	-	-	-
	White-footed Mouse	22	alluvial stream bottoms; open woodlands; wooded slopes/slope bases;	-	-	-	-
	Fulvous Harvest Mouse	18	open woodlands; alluvial stream bottoms; rare in prairies/open fields;	-	-	-	-
	Eastern Harvest Mouse	12.5	prairies/open fields; disturbed grounds; thickets;	-	-	-	-
	Plains Harvest Mouse	8	prairies/open fields;	-	-	-	-
	Northern Pygmy Mouse	8	prairies/open fields; open woodlands; disturbed grounds;	-	-	-	-
	Desert Shrew	6.5	xeric habitats;	-	-	-	-
	Least Shrew	5.75	prairies/open fields;	1078	-	1 kg	Simms 1984

maps presented by Davis and Schmidly (1994), and exclude bats. The occurrence of members in the *Ursidae* (bear), *Antilocapridae* (antelope), and Bison (buffalo) families is based on former distributions reconstructed by Weniger (1997). The distribution of the members in the remaining families is based on present distributions.

Table 8-2 groups the 51 species inhabiting the Gulf Plains and Marshes into body size categories. It indicates that as body size decreases the number of species in each size category increases, with the exception of small (25) to very small (20). Three species are grouped into the large body size category: bison, black bear, and mountain lion. Seven species are in the medium body size category (e.g., white-tailed deer, pronghorn, wolf, collared peccary (javelina), 21 in the small category (e.g., squirrels, skunks, opossum, rabbits), and 20 species are in the very small body size category (e.g., gophers, rats, mice).

Nutritional data of interest to us was not found in the available literature for the species living along the Texas Coast, though some data is reported for the same species living elsewhere. Likewise, no data is reported here on seasonal fluctuations. Reidhead (1976) does report nutritional data on species in Indiana (mostly from Watt and Merrill 1963) and uses fat grams reported there to determine seasonal fluctuations of protein which is then used to determine other micro-nutrient values mathematically.

Deer was likely the principal medium body size mammal within the coastal plains. As such, it is worth discussing in some detail the possible population dynamics that may have influenced availability and nutritional condition and therefore procurement practices and reliance on alternative resources. Statistics provided by the Texas Parks and Wildlife Department indicate that the Gulf Coastal Plains and Marshes vegetation region of the state is ranked second among the ten vegetation regions of the state in terms of the number of deer killed annually per 1,000 acres, with 6.3 deer/1,000 acres harvested per year. The number of deer killed by hunters is dependent on several factors, including the number of hunters. The number of hunters, however, is controlled at the state level through the issuance of licenses with the number of licenses issued dependent on densities of deer within these broad vegetation regions. The number of kills, then,

Table 8-2. Breakdown of Mammalian Species by Body Size Category

Size Category	Totals
Large (> 100,000 g.)	3
Medium (53,000-10,000 g.)	7
Small (8,500-400 g.)	21
Very Small (275-5 g.)	20
Totals	51

(Compiled from Davis and Schmidly 1994)

may roughly approximate relative trends in deer populations within a region.

The food supply found within a deer's home range tends to be the principal factor limiting deer population sizes (Leopold et al. 1947; Walsh 1985:115). According to Hiller (1996:19) forbs, such as weeds and broadleaf flowering plants, are the preferred and principal food (60-80% of diet) of whitetail deer in the Coastal Bend. Browse, the tender leaves and young shoots of woody plants, tends to also be important while the tender green growth of cool-season grasses represents only roughly 15% of the mid-to late-winter diet. Armstrong and Young (2001:11) have noted the critical importance of browse, especially the mast (nuts) production during the fall, because it provides the key fat reserves that allow deer to over-winter without starving. Given the winter nutritional bottleneck and the role of sufficient high quality browse and mast crop for deer survival, the annual number of deer killed may be roughly correlated with total deer populations within each region, and this in turn is conditioned by the amount of total browse and mast production the year before within each vegetation region.

This observation seems to be supported by estimated whitetail deer populations in the Gulf Coastal Plains compared to the South Texas Plains. Between 1974 and 1978, it has been estimated that the coastal plains supported a mean total deer population of 81,309 deer. In contrast, it was estimated that within the South Texas Plains and Brush Country, the mean deer population was 652,566 during the same period (Liebow et al. 1980:242), an eight-fold difference. Present estimates of deer population densities also support the regional differences in deer densities. The Quality Deer Management Association's Whitetail Map Guide (2008) shows that on the Edwards Plateau, deer densities tend to be greater than 45 animals per square mile. In contrast, deer densities drop to around 30-45 animals per square mile on the Coastal Plain and the South Texas Plains and Brush Country has even lower densities (15-30 deer/sq. mile).

The apparent correlation between mast production and deer population densities also suggests that deer population densities will vary to some extent by habitat distribution. That is, deer will more likely be encountered in higher densities in the prairie-riverine interface habitat and in oak motts and stands scattered across the prairie habitat than on the open prairie, especially during the winter. Support for this proposition is found in deer population estimates for a number of the counties in the Gulf Plains and Marshes. For instance, between 1977 and 1978 deer populations of San Patricio and Refugio Counties were estimated at a mean of 90.8 and 50.8 animals per 1,000 acres, respectively. Significantly, in 1976

the total area of each county occupied by fluvial woodlands and oak motts combined was approximately 35.1 mi.² (San Patricio) and 18.0 mi.² (Refugio), respectively (extrapolated, based on figures provided in Brown et al. 1976). In contrast, Aransas County, has a total of approximately 13 mi.² of fluvial woodlands and oak motts, had an estimated mean of 17.1 deer per 1,000 acres (Liebow et al. 1980:240-241; Brown et al. 1976:Table 7).

Studies of deer population behavior on the more open Gulf Plains and Marshes also indicate that during the daytime hours deer tend to spend more time in brush-heavy areas and feed in more open areas only very early or very late in the day when they are under the cover of low light conditions (Walsh 1985). The role of vegetation cover is so critical for deer densities that Steuter and Wright (1980) concluded after studying deer population densities on the Piloncillo Ranch in Dimmit and Webb Counties that as brush cover decreased whitetail deer densities also declined across the ranch. In addition to lower population densities, deer in the Gulf Coastal Plains and Marshes tend to be smaller and therefore weigh less than deer in the Edwards Plateau (D. Forrester, personal communication November 2006). While influenced by deer population management strategies, bucks on the Edwards Plateau are estimated to weigh roughly 135 lb. In contrast, bucks on the Coastal Plains rarely exceed 110 lb. (M. Lockwood, personal communication November 2006; D. Forrester, personal communication November 2006). The final characteristic of deer populations worth highlighting is their seasonal weight change that tends to be particularly noticeable among bucks. Deer managers and hunters have estimated that during the rut, bucks may lose as much as 20-30% of their body weight while does will lose a much lower percentage (e.g., 5-10%; D. Forrester, personal communication November 2006). The timing of breeding season in white tail deer can vary a great deal, depending on environment. Even within Texas, for instance, there is considerable variation in the timing of the rut, from roughly mid-November to the end of December in the Trans Pecos area to September to October in the northern part of the Gulf Prairies and Marshes area (see Texas Parks and Wildlife 2005).

All of the above can inform estimates of deer population change during climate changes of the past, suggesting that during more xeric periods, the deer population in the vicinity of 41NU2 would have been much lower, and, especially during the fall and winter months, largely confined to densely covered areas close to the rivers and creeks, where oak and other mast producing species would still be present. In contrast, during more mesic periods, the number of deer as well as their size and ubiquity would increase, due to the increased cover and increased area covered by mast producing species.

While the open prairie may not be the preferred habitat for deer it would have been prime habitat for pronghorn antelope. Therefore, as suggested by historic distribution data (Weniger 1997:43), it is likely that prior to its extermination from the region, the Gulf Plains and Marshes region supported a viable antelope population. Furthermore, given its habitat preference for wide open country with high visibility, it is likely that antelope densities would have been high exactly in those segments of the prairie habitat that would be less preferred by deer and populations would have been higher in more xeric climates when deer populations would be lower.

The Gulf Coastal Plains and Marshes region is at the hypothetical southern edge of the Plains historic bison distribution (Figure 8-1). Nonetheless, it is likely that the prairie with an abundance of big bluestem and other native tall and mid grasses, dissected by numerous perennial streams, would have offered ideal habitat for bison herds. Therefore, it is necessary to consider their presence in the region as a significant part of any models attempting to reconstruct prehistoric hunter-gatherer adaptations in the region.

In addition to the mammals mentioned above, reptiles and amphibians often can represent an additional source of protein that can be acquired while on foraging trips (Cabeza de Vaca 1555:52; T.N. and T.J. Campbell 1988:20; Marshall

1976). Data from Brown (1950:Table III) indicates that 23 species of amphibians (e.g., salamanders and frogs) and 67 species of reptiles (e.g., lizards, snakes, alligators and turtles) are present within the Post Oak Belt of Texas, a region that is roughly equivalent to the Gulf Coastal Plains. Among the amphibians there are 17 species of frogs, while among the reptiles, snakes (38 species), and turtles (15 species) have the highest species diversity.

Estuary and Freshwater Fish

Bay Estuary-Lagoon Species

A review of the marine fishes that can be found in estuarine settings summarized by Hoese and Moore (1977) indicates that 14 species of fish are regularly present in the bay systems of the Texas coast. They include: black drum (*Pogonias cromis*), spotted seatrout (*Cynoscion nebulosus*), sand seatrout (*Cynoscion arenarius*), sea catfish (*Arius felis*), Gafftopsail catfish (*Bagre marinus*), pinfish (*Lagodon rhomboids*), spot (*Leiostomus xanthurus*), red drum or redfish (*Sciaenops ocellata*), sheepshead (*Arcosargus probatocephalus*), Gulf menhaden (*Brevoortia patronus*), southern and Gulf flounder (*Paralichthys lethostigma*), Atlantic croaker (*Micropogonias undulates*), striped mullet (*Mugil cephalus*) and the bay anchovy (*Anchoa mitchilli*).

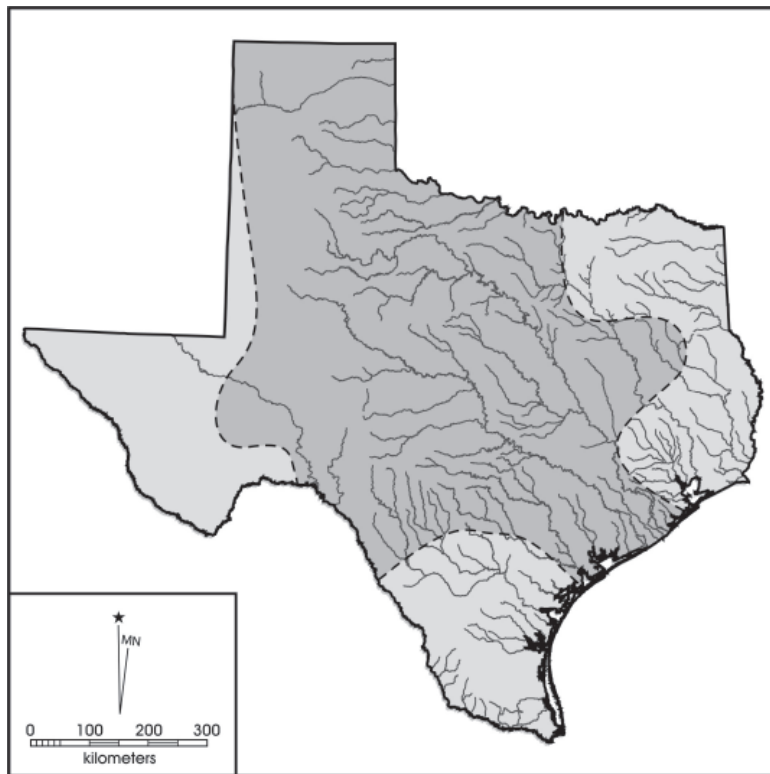


Figure 8-1. Hypothesized historic bison population distribution across Texas (based on data from Weniger 1997).

In addition to the number of economically significant species found within the region, the Laguna Madre and Corpus Christi Bay are locations where the highest densities of Black Drum have been noted in surveys and Matagorda Bay is considered a nursery area for bay anchovy. Interestingly, however, of these 15 species, only spotted seatrout, Atlantic croaker, black drum, red drum and sheepshead have been consistently recovered in archeological sites on the coast (Ricklis 1997; Zimmerman 1997).

Table 8-3 presents a summary of the habitat conditions, specifically salinity and temperature ranges, tolerated and preferred by these species as well as spawning periods and settings and juvenile and adult presence in accessible bay estuary-lagoon settings. The size range of adult fish is also provided for each of the species. They range from 45-970 mm in size. The data have been compiled from Hoese and Moore (1977) and individual species life histories produced by the Coastal Ecology Group Waterways Experimental Station (Collins 1985; Johnson and Seaman 1986; Lassuy 1983a, 1983b; Morton 1989; Reagan 1985; Reagan and Wingo 1985; Sutter and McIlwaine 1987; Sutter et al. 1986).

Table 8-3. List of Marine Fish that Regularly Invade Gulf Coast Freshwater Habitats

Species	Salinity (ppt)		Temperature (°C)		Spawning Period	Age Group in Estuary	When in Estuary	Adult Size Range (mm)	Additional information	Calories in 100 g raw	Reference
	Tolerates	Pref.	Tolerates	Pref.							
Atlantic Croaker	0-70	15-19	3-36	13-28	Sept-Mar; Nov peak	Juveniles and adults	Apr-Aug	195-610	Lassuy 1983a	96	Watt and Merrill 1963
Bay Anchovy	0-45	9-35	5-40	12-27	Mar-May	Juveniles and adults	All year	45-100	Robinette 1983	–	–
Black Drum	0-80	9-26	2.7-36	12-33	Feb-Mar;	Juveniles and adults	Juveniles remain until maturity. Adults April-Feb	285-990	Sutter et al. 1986	80	Watt and Merrill 1963
Gafftopsail Catfish	0-40	5-30	20-35	25-37	May-Aug	Juvenile and adult	Feb-Nov	150-571	Muncy and Wingo 1983	–	–
Gulf Flounder	6-35 ^a	22-35	8.3-32.5	>23	Oct-Dec	Juveniles and adults	Juveniles remain until mature; Adults Mar-Oct	100-500	Gilbert 1986	–	–
Gulf Menhaden	0-75	20-60	5-35	23-30	Oct-Mar	Juvenile to adult	Mar-Sept	255-455	Lassuy 1983b	–	–
Pinfish	0-75	0-35	10-36	18-23	Sept-Dec, Oct peak	Juvenile to adult (except during	Juveniles remain until mature; Adults: Jan-Aug	115-240	Muncy 1984	–	–
Red Drum	0-50	10-20	2-37	13-28	Aug-Dec	Juveniles, young adults	Juveniles may remain all year, young adults Jan-July	283-970	Reagan 1985	80	Watt and Merrill 1963
Sand Seatrout	0-45	0-35	5-35	20-30	Mar-May; Aug-Sept	Juveniles and adults	All year, though some adults spawn in Gulf	140-342	Sutter and McIlwain 1987	–	–
Sea Catfish	0-40	30-40	5-37	25-37	May-Aug	Juvenile and adult	Feb-Nov	120-495	Muncy and Wingo 1983	–	–
Sheephead	0-35	5-30	5-35	8-29	Mar-May	Juveniles and adults	Jun-Nov	200-720	Jennings 1985	113	Watt and Merrill 1963
Southern Flounder	0-36	15-35	5-35	14-22	Oct-Dec	Juveniles and adults	Juveniles Feb-Nov; Adults Apr-Nov	84-480	Reagan and Wingo 1985	–	–
Spot	0-66	–	6-36.7	10-30	Dec-Mar	Juveniles and adults	Apr-Dec	130-210	Hales and Van Den Avyle 1989	219	Watt and Merrill 1963
Spotted Seatrout	0-77	15-35	5-35	15-27	Mar-Sept	Juvenile and adults	Juveniles remain until maturity. Adults move in and out of estuary throughout the year depending on temperature and salinity.	–	Johnson and Seaman 1986	–	–
Striped Mullet	0-75	0-35	10-33	16-30	Oct-Feb; Nov-Dec peak	Juveniles and adults	Juveniles remain until maturity. Adults Mar-Sept	195-300	Collins 1985	146	Watt and Merrill 1963

Marine fishes that range into the Texas estuaries must tolerate a wide range of temperature and salinity conditions. Striped Mullet is considered to be more tolerant of temperature changes than many other species. In contrast, the Gulf menhaden is less tolerant of rapid changes in temperature and salinity than other species and the black drum and bay anchovies have narrower tolerance ranges of salinity and

water temperature conditions. This is exemplified by the fact that when hypersaline conditions prevailed in the Laguna Madre during the severe drought in 1953 some black drum moved upstream into tributary streams of Baffin Bay. During this period, fish moved into Corpus Christi, Aransas and San Antonio Bays and black drum moved upstream into Oso Creek to escape from hypersaline conditions in the upper

Laguna Madre (Simmons and Breuer 1982:69). Episodes of mass mortality have been reported among Gulf menhaden under hypersaline conditions reaching 80 ppt or higher (Simmons 1957).

The same response is elicited by more sudden, less predictable changes in ambient temperature as well. For instance, red drum and spotted seatrout move to deeper waters within bays or offshore during blizzards. The consequences of not relocating during sudden cold spells can be disastrous as in 1940 when a large number of spotted seatrout were killed by a cold wave in January (Gunter 1941). Temperatures fell to 39°F on the open bay and thousands of fish were killed. A similarly massive fish kill was recorded in 1951 (Gunter and Hildebrand 1951). That year a cold wave that hit the coast between the January 28 and the February 3 resulted in a drop in open bay temperatures to 38°F and ice had formed along large portions of the shore. Speckled trout, sheepshead, drum, croakers, pin fish, shad and anchovies were among the thousands of dead fish. The Game Fish and Oyster Commission reported that between 60 and 90 million pounds of fish were killed in 1951 (Gunter and Hildebrand 1951). Hurricanes are another source of fish movements (Hoese 1960).

In the majority of the species that use the estuaries as nurseries for juvenile fish, spawning tends to occur in the late summer through winter although sand and silver seatrout may have a summer spawning period (Sutter and McIlwain 1987:5) and black drum may spawn into June (Sutter et al. 1986:3). Most species spawn in off-shore settings (Sutter and McIlwain 1987:5; Sutter et al. 1986:3) although spotted seatrout spawn in estuaries (Hoese and Moore 1977:204-205; Johnson and Seaman 1986:4). Estuaries serve as sanctuaries for many other juvenile fish species as well (Hoese and Moore 1977:205; Morton 1989:6; Reagan 1985:4; Sutter et al. 1986:4). While adult Gulf menhaden are rarely present in estuarine settings (Lassuy 1983b:4-5), adult Atlantic croaker (Hoese and Moore 1977:204), black drum (Hoese and Moore 1977:204), red drum (Hoese and Moore 1977:205), spotted seatrout (Hoese and Moore 1977:204-205), and stripped mullet (Collins 1985:3) are present year-round. Adult spot are present in estuaries during summer and fall (Hoese and Moore 1977:206), and southern flounder are in estuarine systems during the winter and well into the summer (Hoese and Moore 1977:258).

Because many species tend to seek deeper, warmer waters as water temperatures decrease during the winter (Gunter 1938; Gelwick et al. 2001), it has been suggested that seasonal changes in ambient and water temperature are responsible for the immigration and emigration of fish to and from estuaries into offshore waters. For instance, juveniles of the Atlantic

croaker return to the estuaries in spring as ambient and water temperatures rise and begin their seaward migration in October-November during the period of greatest average drop in surface water temperature between successive months (Lassuy 1983a:4). Similarly, post larvae southern flounder also move into estuaries in late February followed by adults in April-June as water temperatures rise above 16°C (Reagan and Wingo 1985:3).

In studying seasonal variations in the abundance of estuarine and marine fishes in Barataria Bay on the Gulf of Mexico in Louisiana, Gunter (1938:342) has found that in the case of some species of fish, migrations between more open gulf waters and the bay “ was a recurrent, sequential phase in the life history or annual cycle of the fish...” The life cycle pattern is also likely tied to changes in the seasonal abundance of food and is in turn a product of long-term evolutionary adaptations to a variety of factors (i.e., food availability, salinity, temperature, etc).

Gunter (1938), based on catch statistics with otter trawls in bay and gulf settings, demonstrates that with a few exceptions (i.e., yellowtail [*Bairdiella chrysura*]; spot [*Leiostomus xanthurus*]; spotted trout [*Cynoscion nebulosus*]; menhaden [*Brevoortia patronus*]; anchovies [*Anchoviella epsetus*]; and hardhead [*Galeichthys felis*]), the bulk of the species are not present in large numbers in the bays and estuaries during the winter months between November and February. During these months both adults and small and juvenile fish are found in the deeper waters of the Gulf rather than the shallow waters of the bays and estuaries. In the case of these species, they register peaks in abundance in bays and estuaries from the early spring and through the summer months beginning with the rise in water temperature and salinity and the increase in food availability (Gunter 1938:Figures 2-3, 7, 9-11,14). The croaker (*Micropogon undulatus*) is a good example of this pattern (Gunter 1938:322-324). The adults spawn outside of the bays near passes during late fall and winter and come into them in large schools between February and July. The abundance of croakers peaks between February and March as the very small and small specimens come into the fishery and nursery bays and estuaries.

A number of variations are apparent among the species that tend to be residents of the bays during the winter (i.e., yellowtail, menhaden, and anchovies). For instance, the spot and the spotted trout spawn from late December through March. The spot may spawn either in the bays or the open gulf (Gunter 1938:326-328), while the spotted trout typically spawns in the bays (Gunter 1938:330-331). The bay spawning of the spot and the spotted trout results in large numbers of adult animals present in the bays during the winter (Gunter

1938: Figures 6 and 8). Menhaden and anchovies spawn during the summer in the gulf and the small and juvenile fish form a pronounced peak in abundance in the bays and estuaries during the winter months (Gunter 1938:335-336: Figure 12; 337, Figure 13).

During spawning or the migration of very small to juvenile fish in or out of the bays and estuaries many species form large schools of individuals. This schooling of fish is an important behavior that results in the concentration and en-mass movement of large numbers of fish. Adult spotted sea trout and bay anchovy school as do juvenile red drum and Gulf menhaden. This behavior results in high densities of fish that enhances their efficient procurement even with the most rudimentary technologies (e.g., baskets, trays; B. Ojeda, personal communication 1979, based on observation on the Pacific coast of Peru). In effect, schooling fish represent hyper-productive resources that can be efficiently procured with relatively rudimentary technologies.

Overall, these patterns suggest that contrary to expectations, fish populations are not available year-round and that extreme variability exists in their seasonal occurrence as well as in their abundance and age profile during the year within the bays and estuaries that would be most easily accessible by prehistoric hunter-gatherers. Cabeza de Vaca himself hinted at such seasonal changes in availability when he noted that

while fish were taken with cane wears between October and February, by the later part of this period fish were no longer abundant (see Chapter 7).

Freshwater Fish

Presently, 72 species of freshwater fishes occur in the streams cross-cutting the Coastal Plains and emptying into the Gulf of Mexico (Lee et al. 1980). Of these, 28 species have average standard lengths or total lengths (Table 8-4) of 115 mm or greater, or equivalent to the smallest adult marine fish that may be captured in Corpus Christi Bay. In addition, several of these (i.e., freshwater drum, smallmouth buffalo, black buffalo, bowfin and spotted gar) can get as large as 500 mm or more, well exceeding the maximum size of nearly all marine species that frequent primary and secondary bays. The larger species tend to be limited to the larger permanent streams. Smaller species can survive in intermittent prairie streams where small pools or pockets of standing water can provide protection for enough individuals to repopulate the stream when it is flowing again (Meador et al. 1990). It is expected that some degree of seasonal behavior and variation in abundance is exemplified by freshwater species as well driven by water temperature and life cycle factors. At a minimum, it is expected that during the colder months, fish will tend to migrate to deeper pools of water that will provide some insulation and fish abundances will be higher during the May-June or October rainy season (Chapter 2; Figure 2-5).

Table 8-4. Freshwater Fish Species within the Gulf Coastal Plains and Marshes

Common Name	Scientific Name	Length (mm)	Calories in 100 g raw	Reference
Freshwater Drum	<i>Aplodinotus grunniens</i>	10000 g	121	Watt and Merrill
Spotted Gar	<i>Lepisosteus oculatus</i>	850 SL	–	–
Smallmouth Buffalo	<i>Ictiobus bubalus</i>	585 TL	113	Watt and Merrill
Bowfin	<i>Amia calva</i>	534 TL	–	–
Black Buffalo	<i>Ictiobus niger</i>	521 TL	113	Watt and Merrill
Spotted Bass	<i>Micropterus punctulatus</i>	300 SL	–	–
Largemouth Bass	<i>Micropterus salmoides</i>	410 TL	104	Watt and Merrill
Longnose Gar	<i>Lepisosteus osseus</i>	360 TL	–	–
White Bass	<i>Morone chrysops</i>	348 TL	98	Watt and Merrill
River Carpsucker	<i>Carpionodes carpio</i>	343 TL	–	–
Gizzard Shad	<i>Dorosoma cepedianum</i>	288 TL	200	Watt and Merrill
West Mexican Redhorse	<i>Moxostoma congestum</i>	278 SL	98	Watt and Merrill
Black Crappie	<i>Pomoxis nigromaculatus</i>	275 TL	79	Watt and Merrill
Blue Sucker	<i>Cylocheilus elongatus</i>	245 SL	104	Watt and Merrill
Silver Chub	<i>Hybopsis storeriana</i>	231 TL	145	Watt and Merrill
Green Sunfish	<i>Lepomis cyanellus</i>	200 TL	–	–
Redear Sunfish	<i>Lepomis microlophus</i>	192 TL	–	–

SL=standard length, measured from tip of snout to posterior end of caudal skeleton, does not include tail fin.

TL=total length, measured from tip of snout to posterior end of tail fin.

SL was used when possible, lengths are averages.

Note: length not given for freshwater drum - average weight in grams is given instead. 83 species total.

Table 8-4. Continued...

Common Name	Scientific Name	Length (mm)	Calories in 100 g raw	Reference
Bluegill	<i>Lepomis macrochirus</i>	191 TL	–	–
Striped Mullet	<i>Mugil cephalus</i>	190 SL	146	Watt and Merrill
Blue Catfish	<i>Ictalurus furcatus</i>	170 SL	103	Watt and Merrill
Warmouth	<i>Lepomis gulosus</i>	168 TL	–	–
White Crappie	<i>Pomoxis annularis</i>	165 SL	79	Watt and Merrill
Common Carp	<i>Cyprinus carpio</i>	151 SL	115	Watt and Merrill
American Eel	<i>Anguilla rostrata</i>	138 TL	233	Watt and Merrill
Stoneroller	<i>Campostoma anomalum</i>	133 SL	–	–
Channel Catfish	<i>Ictalurus punctatus</i>	127 SL	103	Watt and Merrill
Flathead Catfish	<i>Pylodictis olivaris</i>	122 SL	103	Watt and Merrill
Longear Cunfish	<i>Lepomis megalotis</i>	121 SL	–	–
Threadfin Chad	<i>Dorosoma petense</i>	110 SL	170	Watt and Merrill
Golden Chiner	<i>Notemigonus crysoleucas</i>	101 SL	–	–
Black Bullhead	<i>Ictalurus melas</i>	99 SL	72	Watt and Merrill
Spotted Sunfish	<i>Lepomis punctatus</i>	98 SL	–	–
Guadalupe Bass	<i>Micropterus treculi</i>	93 SL	–	–
Mexican Tetra	<i>Astyanax mexicanus</i>	79 SL	–	–
Plains Killfish	<i>Fundulus zebrinus</i>	69 TL	–	–
Yellow Bullhead	<i>Ictalurus natalis</i>	64 SL	84	Watt and Merrill
Central Silvery Minnow	<i>Hybognathus nuchalis</i>	63 SL	–	–
River Darter	<i>Percina shumardi</i>	63 SL	–	–
Blackspot Shiner	<i>Notropis atrocaudalis</i>	62 SL	–	–
Bigscale Logperch	<i>Percina macrolepada</i>	62 SL	–	–
Blackstripe Topminnow	<i>Fundulus notatus</i>	60 TL	–	–
Blacktail Shiner	<i>Notropis venustus</i>	60 SL	–	–
Sailfin Molly	<i>Poecilia latipinna</i>	60 SL	–	–
Weed Shiner	<i>Notropis texanus</i>	58 SL	–	–
Sand Shiner	<i>Notropis stramineus</i>	57 SL	–	–
Texas Shiner	<i>Notropis amabilis</i>	55 SL	–	–
Dusky Darter	<i>Percina sciera</i>	55 SL	–	–
Bantam Sunfish	<i>Lepomis symmetricus</i>	54 SL	–	–
Orangethroat Darter	<i>Etheostoma spectabile</i>	54 SL	–	–
Roundnose Minnow	<i>Dionda apiscopa</i>	52 SL	–	–
Ironcolor Shiner	<i>Notropis chalybaeus</i>	50 TL	–	–
Pallid Shiner	<i>Notropis amnis</i>	50 SL	–	–
Beautiful Shiner	<i>Notropis formosus</i>	49 SL	–	–
Shaprnose Shiner	<i>Notropis oxyrhynchus</i>	48 SL	–	–
Freckled Madtom	<i>Noturus nocturnus</i>	48 SL	–	–
Ribbon Shiner	<i>Notropis fumeus</i>	47 SL	–	–
Silverband Shiner	<i>Notropis shumardi</i>	47 SL	–	–
Red Shiner	<i>Notropis lutrensis</i>	46 SL	–	–
Bluntnose Darter	<i>Etheostoma chlorosomum</i>	46 SL	–	–

SL=standard length, measured from tip of snout to posterior end of caudal skeleton, does not include tail fin.

TL=total length, measured from tip of snout to posterior end of tail fin.

SL was used when possible, lengths are averages.

Note: length not given for freshwater drum - average weight in grams is given instead. 83 species total.

Table 8-4. Continued...

Common Name	Scientific Name	Length (mm)	Calories in 100 g raw	Reference
Chub Shiner	<i>Notropis potteri</i>	45 SL	–	–
Tadpole Madtom	<i>Noturus gyrinus</i>	44 SL	–	–
Sheephead Minnow	<i>Cyprinodon variegatus</i>	43 SL	–	–
Bullhead Minnow	<i>Pimephales vigilax</i>	42 SL	–	–
Pugnose Minnow	<i>Notropis emiliae</i>	41 SL	–	–
Mimic Shiner	<i>Notropis volucellus</i>	37 SL	–	–
Rainwater Killfish	<i>Lucania parva</i>	37 SL	–	–
Ghost Shiner	<i>Notropis buchanani</i>	35 SL	–	–
Speckled Chub	<i>Hybopsis aestivalis</i>	34 SL	145	Watt and Merrill
Slough Darter	<i>Etheostoma gracile</i>	33 SL	–	–
Cypress Darter	<i>Etheostoma proeliare</i>	33 SL	–	–
Backwater Darter	<i>Etheostoma zoniferum</i>	32 SL	–	–
Greenthroat Darter	<i>Etheostoma lepidum</i>	32 SL	–	–
Mosquitofish	<i>Gambusia affinis</i>	28 TL	–	–

SL=standard length, measured from tip of snout to posterior end of caudal skeleton, does not include tail fin.

TL=total length, measured from tip of snout to posterior end of tail fin.

SL was used when possible, lengths are averages.

Note: length not given for freshwater drum - average weight in grams is given instead. 83 species total.

Estuary and Freshwater Molluscs

The role of shellfish in prehistoric diets has received much attention since the 1970s and gone from marginal resource status to important source of protein (Erlandson 1988), calories, carbohydrates (Claassen (1986), or other mineral nutrients supplementing terrestrial resources during lean months. Though the prehistoric groups consuming these items likely were not cognizant of the specific nutrition they were or were not getting from shellfish, they were likely aware of their fitness levels and how diet effected it.

The nutritional data that was gathered through a search of the available literature on individual species was in no case derived from species collected from the Gulf of Mexico. This fact shows the lack of quantified nutritional data on this as well as the bulk of potential resources that may have been used by prehistoric hunter-gatherers of the region. Much of the nutritional information we present in subsequent tables was compiled from Watt and Merrill (1963) and we can only assume that since the data refers to many of the same species as those found in the Gulf, the information is roughly applicable.

Freshwater Mussels are sessile filter feeders that burrow into the substratum with a large muscular foot that is also used for locomotion. Generally, though, as most bivalves, these spend most of their lives in a small area Freshwater mussels are quite long-lived, reaching 10-80 years. Hard clams inhabit subtidal regions of bays and estuaries to 15 m deep in firm bottom mud flats or areas of sand and shell. They can tolerate

a wide range of salinities. Oysters are usually found in groups as oyster reefs formed by oysters furrowing on a foundation of dead shells. Clusters join to form large continuous groups. They also provide habitats for other marine resources. Scallops prefer shallow marine habitats with lots of sea grass. These can move considerable distances within and between habitats unlike other bivalves (Pearse et al. 1987).

We do not account for the size or seasonal fluctuation of the species though we acknowledge that these are critical to the nutritional content of the individual. Accordingly, Parmalee and Klippel (1974) collected freshwater mussels in the Midwest to report the nutritional value of shellfish. Of relevance here is the wide range of size, weight, and meat to shell ratio between and within freshwater mussels that certainly would affect nutritional value. Their study included the range and mean weights and measurements of mussel species collected. While only seven of the species collected from the Mississippi River area are also in the present study region, we can assume that variability in specimen sizes will be found in our study area.

Shellfish are general a poor source of calories and pursuit of mollusks to meet caloric needs would likely result in over-harvesting. Most groups will likely get their calories from plants and other animals. Those relying on carbohydrates for calories would need a balanced amount of protein that shellfish in part could provide under optimal conditions (Erlandson 1988). Case studies in areas other than the gulf coast have shown minimum daily requirements of protein

can be collected from shellfish in a relative short time (D.W. Bird 1996), particularly in spring when carbohydrate and protein content may peak prior to the reproductive season (Erlandson 1988).

Bay-Estuary-Lagoon Molluscs

The list of bay-estuary-lagoon system molluscs provided in Table 8-5 was compiled based on Andrews (1977) and is limited to species with adult maximum length dimensions larger than 35 mm. It is assumed that while smaller species may turn up in archeological assemblages it is most likely that they were incidentally recovered due to harvesting methods employed rather than purposefully harvested for food. We assume that the amount of meat present in such small species would provide very low returns making the systematic harvest of such small species extremely inefficient. The species of molluscs noted in archeological assemblages (Ricklis 1997; Zimmerman 1997) are identified by a checkmark.

Seventy-three species are listed in the table with the top 38 on the list being species that are regularly found in bay-estuary-lagoon settings and would be readily accessible to hunter-gatherers. The bottom 35 species inhabit offshore environments and may be less accessible to collectors. The 73 species that fit the size cutoff compares well with results of systematic sampling of Corpus Christi Bay between 1981 and 1982 (Castiglione 1983). That survey yielded a total of 165 species of molluscs of which 89 species were gastropods and 76 were pelecypods, also called bivalves (Castiglione 1983). Similarly, a three-year survey (1972-1975) of molluscs in Corpus Christi Bay, Copano Bay and Aransas Bay identified 68 species of pelecypods (Holland et al. 1973, 1974, and 1975). The Bureau of Economic Geology (BEG) sponsored a similar survey of Copano Bay between 1976 and 1977; it identified 33 species of pelecypods (Calnan 1980). The smaller number of species is most likely reflective of the single bay system surveyed. In terms of mollusc densities, Castiglione (1983:52) found that the average total number of individuals for six of the more common species found in samples from three distinct localities ranged from 427 to 992 specimens per meter square (April 1991 to March 1992). A total mean number of 1988/m² molluscs of these six species were recovered at the three seagrass-dominated sample localities during the year.

In general, sediment type (i.e., substrate) and diversity, organic content, and predation coupled with levels of salinity strongly influence the species composition and density of mollusc assemblages encountered at any one time and in any one place within coastal bay-estuary-lagoon habitat (Holland et al. 1975). For instance, Castiglione (1983) found

that portions of Corpus Christi Bay with high proportions of sand in the substrate had low mollusc productivity (yearly mean of 150 molluscs/m²) but high species diversity. In contrast, a sample area with 90% mud substrate had both low productivity (yearly mean of 28 molluscs/m²) and diversity. Areas with mixed substrate of shell, sand and mud had low species diversity but high productivity with a small number of species dominating mollusc assemblages (Castiglione 1983:90-91).

As with other resources already discussed, seasonal variation in availability and nutritional condition is also present among estuarine mollusks. Castiglione (1983), for instance, found that the highest population densities occur during the months of February-April. Carbohydrate and protein content may peak during spring prior to the reproductive season. Mollusc population densities decreased during the summer and fall (Castiglione 1983:91). An influx of young individuals during the spring followed by a peak in mortality rates accounts for the trend. These data support Texas coastal marine subsistence studies that find a reliance of fish and shellfish during winter months (Ricklis 1996). Shellfish gathering would also be more efficient when tides are lowest during daylight hours, which occur during winter.

In addition to seasonal variability in abundance, levels of salinity and water temperature significantly influence the residence time and degree of mobility of mollusc populations within and between estuary and offshore settings. In addition to preferred substrate/habitat information for each common species, Table 8-5 also presents the salinity tolerance limits and preferred substrate for the principal mollusc species of potential economic utility to human inhabitants of the coast. The majority can tolerate a wide range of salinity conditions and most importantly can live and reproduce in high-salinity environments. Only three bay-estuarine-lagoon bivalve species (*Polymesoda caroliniana*; *Rangia cuneata*, and *R. flexuosa*) are adapted to low (5-19 ppt) salinity. Twenty-three species tolerate a broad range (15-40 ppt) of salinity values although their preferred range is in the lower half, while 48 species are adapted to a narrower, higher range of salinity (25-40 ppt) preferring habitats with salinity values in the upper half of this range.

While the differences in tolerated and preferred salinity levels are critical in understanding the settings where particular species may be harvested, they are also important as barometers of changes in habitat conditions over time or in space (Parker 1955). Studies of changes in mollusc assemblages of coastal bays and estuaries during significant short and long-term weather events or climatic changes have demonstrated that mollusc assemblages are quite vulnerable

Table 8-5. Common Molluscs Species Present in the Bay-Estuary-Lagoon Habitat

Common Name	Scientific Name	Size Range (mm)	Substrate/Habitat	Salinity	Locality	Type	Calories (C)	Protein (g)	Fat (g)	Carbs (g)	Reference
Atlantic Bay Scallop	<i>Argopecten irradians</i>	50.8-76	bays and open lagoons	15-40%	entire	Scallops	81	15.3	0.2	3.3	Watt and Merrill :56 item 2023
Eastern Oyster	<i>Crassostrea virginica</i>	51-150	brackish bays and estuaries	15-40%	entire	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Fallen Angel Wings	<i>Barnea truncata</i>	71	clay bottoms of bays	15-40%	central	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Texas Venus	<i>Allochorda texaniana</i>	38-76	clay lagoon centers and along clay shores	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Giant Atlantic Cockle	<i>Laevicardium robustum</i>	76-101	close-to-shore and inlet-influenced areas	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Stout Tagelus	<i>Tagelus plebeius</i>	50-68	enclosed lagoons, bay margins	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Carolina Marsh Clam	<i>Polymesoda caroliniana</i>	25-38	estuaries	5-19%	Matagorda Bay and east	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Thick Lucina	<i>Lucina pectinata</i>	25-63	grass flats in open-bay margins and hypersaline lagoons	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Horse or Crested Oyster	<i>Ostrea equestris</i>	25-51	high-salinity oyster reefs	15-40%	entire	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Common Jingle Shell	<i>Anomia simplex</i>	25-51	hypersaline oyster or rock reef	15-40%	entire	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Flat Tree Oyster	<i>Isognomon alatus</i>	50.8-76	in crevices on jetties	25-40%	southern	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Mahogany Date Mussel	<i>Lithophaga biscuita</i>	25-38	in rocks in both shallow and deep water	25-40%	entire	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Scissor Date Mussel	<i>Lithophaga aristata</i>	25-38	in rocks in both shallow and deep water	25-40%	central and south	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Atlantic Cyclinella	<i>Cyclinella tenuis</i>	25-50	inlet-influenced areas and bay margins	25-40%	entire more south	-	-	-	-	-	-
Yellow Cockle	<i>Trachycardium muricatum</i>	51	inlet-influenced areas, bay margins	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Buttercup Lucina	<i>Cyclinella tenuis</i>	38-51	inlet-influenced areas, bay margins, hypersaline lagoons	25-40%	entire	-	-	-	-	-	-
White Crested Tellin	<i>Tellidora cristata</i>	up to 37	inlets and channels	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Green Jackknife Clam	<i>Solen viridis</i>	up to 51	inlets, near shore	15-40%	east	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Transverse Ark	<i>Anadara transversa</i>	12-36	litoral to 10.8 m	25-40%	entire	Clams	49	6.5	0.4	-	-
White Bearded Ark	<i>Barbatia candida</i>	25.4-50.8	on rocks beyond low tide	25-40%	Port Aransas south	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Rough Jingle Shell	<i>Popodesmus rudis</i>	up to 51	on rocks in inlet influenced areas	25-40%	southern	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Mossy Ark	<i>Arca imbricata</i>	25-38	on rocks or firm substratum	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Scorched Mussel	<i>Brachidontes exustus</i>	18-38	on rocks or oyster reefs in bays	15-40%	entire	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Turkey Wing	<i>Arca zebra</i>	50.8	on rocks, shell, roots	25-40%	southern	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Paper Shell	<i>Amygdalum papyria</i>	25-38	open bays and inlet influenced areas	15-40%	entire	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Texas Quahog	<i>Mercenaria campechinesis texana</i>	76-127	open bays and inlet influenced areas	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Purplish Tagelus	<i>Tagelus divisus</i>	25-38	open lagoon margins	15-40%	entire	-	-	-	-	-	-

Table 8-5. Continued...

Common Name	Scientific Name	Size Range (mm)	Substrate/Habitat	Salinity	Locality	Type	Calories (C)	Protein (g)	Fat (g)	Carbs (g)	Reference
Fragile Atlantic Mectra	<i>Maetra fragilis</i>	38-63	open-bay margins	15-40%	southern	Clams	49	6.5	0.4	4.2	–
Constricted Macoma	<i>Macoma constricta</i>	25-63	open-bay margins and centers	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Angel Wing	<i>Cyrtopleura costata</i>	177	open-bay margins inlet infl-areas;	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
False Angel Wing	<i>Petricola pholadiformis</i>	up to 50	open-bay margins, inlet infl areas	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Channel Duck Clam	<i>Raeta plicatella</i>	50-76	outer surf	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Hooked Mussel	<i>Ischadium recurvum</i>	25.6-63.5	oyster reefs in low-salinity bays	15-40%	entire	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Common Rangia	<i>Rangia cuneata</i>	25-63	river-influenced areas	5-19%	central	clams	49	6.5	0.4	4.2	Watt and Merrill 1963:40 item 1363
Brown Rangia	<i>Rangia flexuosa</i>	25-38	river-influenced areas	5-19%	entire, more east	clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Ribbed Mussel	<i>Geukensia demissa granosissima</i>	50.8-101	salt marshes	15-40%	east	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Alternate Tellin	<i>Tellina alternata</i>	70	sand near shore, bay margins, inlets	25-40%	entire more east	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Tiger Lucina	<i>Codakia orbicularis</i>	63-87	shallow flats in sand	25-40%	southern	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Sunray Venus	<i>Callista nimbosa</i>	101-127	shallow sandy bottoms, inlet-infl areas	n/a	entire more south	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Red Brown Ark	<i>Barbatia cancellaria</i>	25.4-38.1	offshore	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Doc Bales' Ark	<i>Barbatia tenera</i>	25-38	offshore	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Saw-Tooth Pen Shell	<i>Atrina serrata</i>	up to 250	offshore	25-40%	entire	–	–	–	–	–	–
Atlantic Pearl Oyster	<i>Pinctada imbricata</i>	38-76	offshore	25-40%	southern	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Calico Scallop	<i>Argopecten gibbus</i>	25-50.8	offshore	25-40%	entire	Scallops	81	15.3	0.2	3.3	Watt and Merrill :56 item 2023
Lion's Paw	<i>Lyropecten nodusus</i>	76-152	offshore	25-40%	southern	Scallops	81	15.3	0.2	3.3	Watt and Merrill :56 item 2023
Ravenel's Scallop	<i>Pecten raveneli</i>	25-51	offshore	25-40%	southern	Scallops	81	15.3	0.2	3.3	Watt and Merrill :56 item 2023
Atlantic Thorny Oyster	<i>Spondylus americanus</i>	76-101	offshore	25-40%	southern	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Spiny Lima	<i>Lima lima</i>	25-51	offshore	25-40%	southern	–	–	–	–	–	–
Atlantic Left-Handed Jewel Box	<i>Pseudochama radians</i>	25-76	offshore	25-40%	entire	–	–	–	–	–	–
Prickly Cockle	<i>Trachycardium isocardia</i>	76	offshore	25-40%	central and south	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Common Egg Cockle	<i>Laevicardium laevigatum</i>	25-51	offshore	25-40%	central and south	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Atlantic Sanguin	<i>Sanguinolaria sanguinolenta</i>	38-50	offshore	25-40%	central and south	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Calico Clam	<i>Callista maculata</i>	38-63	offshore	25-40%	southern	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Elegant Dosinia	<i>Dosinia elegans</i>	50-76	offshore	25-40%	southern	Clams	49	6.5	0.4	4.2	–
Clench's Chione	<i>Chione clenchi</i>	25-63	offshore	25-40%	entire more south	Clams	49	6.5	0.4	4.2	–
Southern Quahog	<i>Mercenaria campechinesis</i>	76-152	offshore	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770

Table 8-5. Continued...

Common Name	Scientific Name	Size Range (mm)	Substrate/Habitat	Salinity	Locality	Type	Calories (C)	Protein (g)	Fat (g)	Carbs (g)	Reference
Geoduck	<i>Panopea bitruncata</i>	127-228	offshore	25-40%	southern	Clams*	–	–	–	–	–
Campeche Angel Wing	<i>Pholas camechinensis</i>	up to 128	offshore	25-40%	entire	Clams	49	6.5	0.4	4.2	–
Half-Naked Pen Shell	<i>Atrina seminuda</i>	235-250	offshore and inlet influenced bays	15-40%	entire	–	–	–	–	–	–
Purplish Semele	<i>Semele purpurascens</i>	25-38	offshore and sandy bottoms and banks	25-40%	southern	Clams	49	6.5	0.4	4.2	–
Atlantic Winged Oyster	<i>Pteria colymbus</i>	38-76	offshore reefs	25-40%	southern	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Tagelus-Like Macoma	<i>Macoma tageliformis</i>	up to 63	offshore, 3.6-19.8 meters	25-40%	entire	Clams	49	6.5	0.4	4.2	–
Taylor's Tellin	<i>Tellina alternata tayloriana</i>	63	offshore, 3.6-21.6 meters; inlet areas	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Pulley's Macoma	<i>Macoma pulleyi</i>	42-55	offshore, delta-influenced waters	25-40%	east	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Lady-in-Waiting Venus	<i>Chione intapurpurea</i>	25-38	offshore; 3.6-21.6 m	25-40%	entire	Clams	49	6.5	0.4	4.2	–
Atlantic Surf Clam	<i>Apisula solidissima</i>	101-127	offshore; 3.6-21.6 meters	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Smooth Duck Clam	<i>Anatina anatina</i>	50-76	offshore; 3.6-21.6 meters	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Disk Dosinia	<i>Dosinia discus</i>	50-76	offshore; 3.6-21.6 meters	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Tulip Mussel	<i>Modiolus americanus</i>	25-101	offshore; below low water mark	25-40%	entire	Mussel	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Leafy Jewel Box	<i>Chama macerophylla</i>	25-76	offshore; deeper calcareous banks	25-40%	southern	Oyster-eastern	66	8.4	1.8	3.4	Watt and Merrill 1963:42 item 1443
Cut-Ribbed Ark	<i>Anadara floridana</i>	63.5-127	offshore; deeper water	25-40%	central and south	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Blood Ark	<i>Anadara ovalis</i>	38-59	offshore; just	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Ponderous Ark	<i>Noetia ponderosa</i>	50.8-63.5	offshore; sandy bottoms	15-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770
Incongruous Ark	<i>Anadara brasiliiana</i>	25-38	offshore; shallow water	25-40%	entire	Clams	49	6.5	0.4	4.2	Watt and Merrill 1963:25 item 770

to extreme fluctuations in salinity and temperature (Hoese 1960). Shallow bays such as Oso Bay and Corpus Christi Bay are especially vulnerable to long-term and short-term changes in climatic condition since the salinity levels and temperature of shallow bodies of water can be changed more rapidly and dramatically even during short climatic perturbations (e.g., cold fronts and sudden downpours) than deep bodies of water.

Freshwater Molluscs

Twenty-seven species of freshwater mussel occur in small and large streams that empty into bays on the Gulf of Mexico and in ponds and lakes found across the Coastal Plains. (Table 8-6; Howells et al. 1996:14). A few of the 27 species (*Unio* spp.) can inhabit intermittent streams and

temporary ponds (Howells et al. 1996:14). However, mussels primarily inhabit flowing waters with substrates of mud, sand, gravel, and cobbles or any combination of these. In areas of strong currents flow obstructions such as fallen trees or gravel bars provide protected areas for mussels. Within Nueces, San Patricio and Refugio counties oxbow lakes (old channels) provide the most suitable habitat for mussels in addition to the free-flowing portions of the river (Neck and Metcalf 1988). Deep soft silts, deep shifting sands, and bedrock or heavy boulder bottoms are limiting habitats and are generally poorly populated (Howells et al. 1996:15). Occurrence of individuals may be sporadic or dense populations may occur in suitable locations. Densities of 81 specimens/m² have been reported for wild populations (Howells et al. 1996:21). Large numbers of crayfish and a variety of small and juvenile fish are found in association with stable mussel beds.

Table 8-6. Common Freshwater Molluscs Found in the Streams, Lakes and Ponds of the Gulf Coastal Plains and Marshes

Common Name	Scientific Name	Length (mm)	Calories (C)	Protein (g)	Fat (g)	Carbs (g)	Source
Threeidge	<i>Amblema plicata</i>	148	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Giant Floater	<i>Anadonta grandis</i>	170	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Paper Pondshell	<i>Anodonta imbecillis</i>	90	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Rock Pocketbook	<i>Arcidens confragosus</i>	160	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Tampico Pearlymussel	<i>Crytonaias tampicoensis</i>	143	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Round Pearlshell	<i>Glebulia rotunda</i>	95	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Texas Fatmucket	<i>Lampsilis bracteata</i>	90	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Louisiana Fatmucket	<i>Lampsilis hydiana</i>	95	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Yellow Sandshell	<i>Lampsilis teres</i>	145	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
White Heelsplitter	<i>Lasmagona complanata</i>	167*	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Fragile Papershell	<i>Leptodea fragilis</i>	133	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Pond Mussel	<i>Ligumia subrostrata</i>	74	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Washboard	<i>Megaloniais nervosa</i>	215	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Pink Papershell	<i>Potamilus ohioensis</i>	141	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Bleufer	<i>Potamilus purpuratus</i>	178	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Southern Mapleleaf	<i>Quadrula apiculata</i>	118	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Golden Orb	<i>Quadrula aurea</i>	77	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Smooth Pimpleback	<i>Quadrula boustonensis</i>	66	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Texas Pimpleback	<i>Quadrula petrina</i>	87	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
False Spike	<i>Quincuncina mitchelli</i>	132	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Squawfoot	<i>Strophitus undulatus</i>	92	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Lilliput	<i>Toxolasma parvus</i>	30	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Texas Lilliput	<i>Toxolasma texasensis</i>	62	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Pistolgrip	<i>Tritogonia verrucosa</i>	170	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Texas Fawnsfoot	<i>Truncilla macrodon</i>	45	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Tapered Pondhorn	<i>Unio merus declivis</i>	121	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363
Pondhorn	<i>Unio merus tetralasmus</i>	141	66	9.6	1.4	3.1	Watt and Merrill 1963:40 item 1363

* shell length

Little is known about variability in the seasonal availability of freshwater mussels (Howells et al. 1996:21). Freshwater mussels occur at varying depths but rarely below about seven and a half meters except during cold weather when some species move to deeper water for protection. Many mussel species can survive for several months during periods of drought by burrowing into the substrate (Howells et al. 1996:18).

Birds

The mosaic of niches found along the coastal bays and estuaries provides a mix of habitats well suited for roosting, resting, foraging, and breeding for a variety of aquatic bird species. As a result, the riverine settings, estuaries and beaches of the Gulf coast represent some of the most important habitat for migrating North America shorebirds (Senner and Howe 1984). In 2006, 247 species were identified by bird watchers in the Corpus Christi area (Corpus Christi Daily 2006). A bird check

list for the Naval Air Station, Corpus Christi, located on the eastern shore of Oso Bay and Corpus Christi Bay lists 310 bird species that have been identified there over the years (U.S. Navy and the National Fish and Wildlife Foundation 2006).

Other than the gallinaceous birds such as the wild turkey and quail that are most common within the prairie habitat, the most economically important group of birds is waterfowl, specifically the family *Anatidae* (ducks and geese) that are most abundant in the marshes of the beach-fluvial interface and the prairie-riverine habitat. Waterfowl are relatively large bodied water birds that are found in highest frequencies on the margins of marshes, ponds, lakes, streams, and the coastline (Texas Game, Fish, and Oyster Commission 1945; Rappole and Blacklock 1994). There are at least 28 species of ducks and geese that migrate through the Gulf coastal regions each fall and winter, or are permanent residents. Most species eat water plants and grasses in areas of shallow

water, but a few species that eat fish prefer deeper waters. Ducks and geese travel in large flocks with up to tens of thousands of individuals in the case of geese. They usually nest on the ground and the number of eggs ranges from 6 to 18. Many species of ducks and geese can be encouraged to land in specific areas by the use of decoys. In the case of some species, such as geese, these decoys can be as simple as fluttering white pieces of paper or cloth (Texas Game, Fish, and Oyster Commission 1945).

A recent study of avian abundance and richness (Carillo 2000) in the area of the Blind (False) Oso indicates that during the 13-month study period more than 33,700 aquatic birds, 77 species, visited the tidal pond, the mudflats and open bay. The highest species richness as well as number of birds was noted in the open bay (n=70 species, 16,123 birds), followed by the wind-tidal flat (n=49 species, 11,201 birds, and tidal pond (n=46 species, 6,418 birds). Figure 8-2 shows the abundance of all birds during the one-year study period. The peak in species richness and abundance occurred in the spring (February-late March) in response to the presence of spring-migrating birds. Species richness and abundance declined between May and early July with the departure of waterfowl and shorebirds for their breeding grounds. A second peak

in richness and abundance occurred between late July and November as fall-migrating species passed through the region. The abundance of birds in different guilds was variable. For instance, wading birds (e.g., reddish and Snowy egrets) were most abundant between November and December. Waterfowl (e.g., Northern shoveler, gadwall, American wigeon, green-winged teal and ruddy duck) were most abundant between October and January (Carillo 2000:129). Gulls and turns (e.g., laughing gull, herring gull, royal turn, black skimmer) were most abundant between late January and late February with a smaller peak between August and November (Carillo 2000:130). Finally, shorebirds (e.g., dowitchers, peeps, stilt sandpipers) were abundant in July and August and again between January and April.

Withers and Chapman (1993) found a similar pattern in an earlier study of seasonal abundance of shorebirds on an Oso Bay mudflat. They documented 26 species of shorebirds during the study year. They found that the winter and early spring (January-March) was the period of highest densities of birds in the study area but the highest species richness was noted during the fall migration (September to October). Low shorebird abundance and diversity were seen during the summer (Withers and Chapman 1993:385-387).

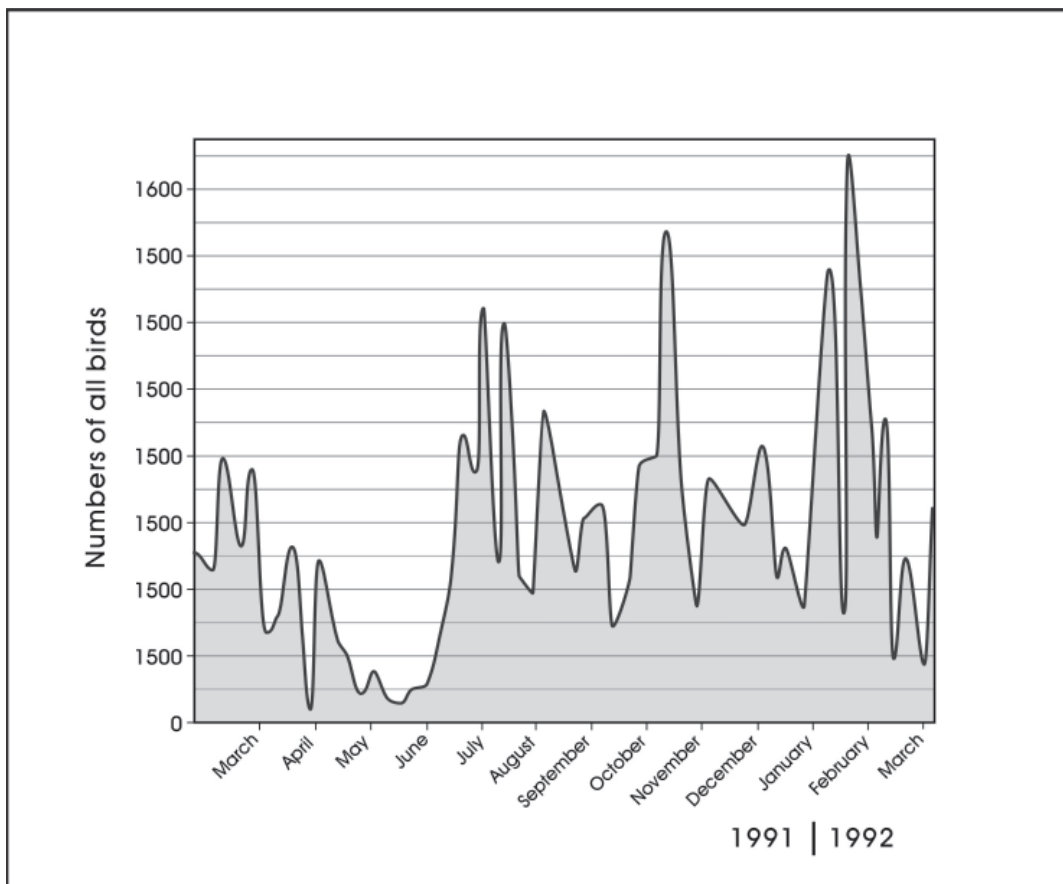


Figure 8-2. Abundance of all birds in the Blind Oso, Oso Bay, Texas (March 1991-March 1992).

While birds offer a ready source of protein and their eggs can provide additional nutrients, the nutritional condition of some species of over-wintering shorebirds is relatively poor and they may not be highly desirable dietary component. For instance, White and Mitchell (1990) have found that the body mass and lipid content of long-billed dowitchers and western sandpipers declined between 44-70% from the time they arrived on the coast in early autumn to their departure in mid-February. The study suggests that at least the members of these species of shorebirds do not acquire substantial fat deposits before their northward migration from the Gulf Coast. However, the reasons for this pattern may be the high overall shorebird densities on the coast at the time and the consequent depletion of food supplies, adverse weather conditions in winter and the possibility that the cost of storing and carrying fat reserves may not exceed the benefits (White and Mitchell 1990:451).

Plant Resources

In general, edible plant resources can be grouped into five major categories depending on the portion of the plant being harvested and consumed. These groups consist of

tubers (including bulbs and roots) fruits, nuts, seeds, and pot-greens. Several publications are available summarizing plant species and their distributions in the state and in South Texas, in particular. The edible plant distribution patterns discussed below are based on reviews of selected ethnobotanical references concerning edible plant species in Texas (Cheatham et al. 1995; Tull 1999), and surrounding states and regions (Ebling 1986; Elliott 1976; Fernald and Kinsey 1943; Kindscher 1987; King 1976; Yarnell 1964).

Figure 8-3 is a graphic representation derived from the above combined references of the number of edible plant species grouped into resource types and broken down by season of availability. The number of plant species utilized as pot-greens is highest during the spring, experiencing a steady drop approaching the fall or autumn months, and declining to a minimum by the winter. Plant species with edible tubers, bulbs, rhizomes, and corms have a bimodal distribution reaching an early peak in the spring, dropping somewhat in the summer, and peaking even higher during the fall before dropping to an annual low during the winter. Plant species with edible fruits tend to peak in the summer and are still relatively common during the fall but occur only in small numbers in

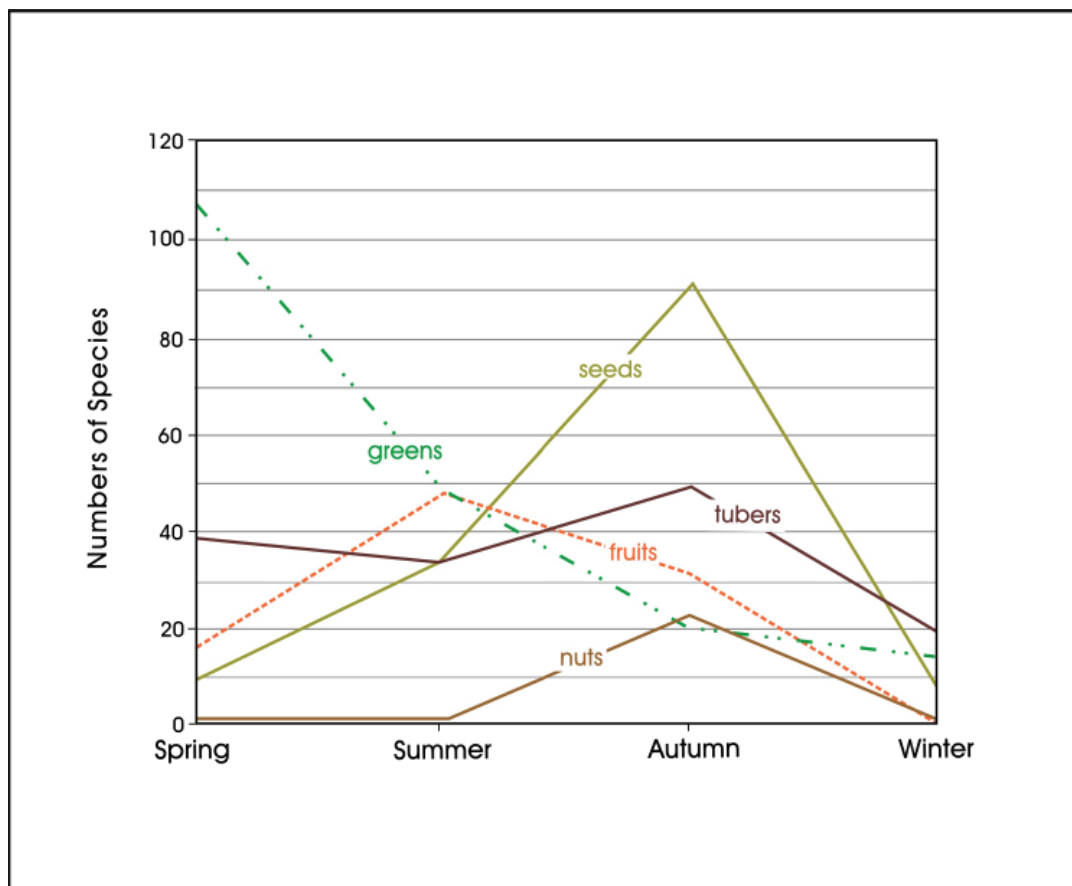


Figure 8-3. Changes in the number of edible plant species by season of availability.

the winter. Species with edible seeds tend to be relatively low in number during the spring, but increase steadily from the late spring, through the summer and peak during the fall. Nuts derived from nut-bearing trees are available primarily during the fall months, peaking in the number of available species between September and October.

Sixty-nine plant species and varieties have been identified as having edible tubers, rhizomes, bulbs, and/or corms. A moderate number of the species are available throughout the year; however, many of these species are apparently more palatable or have higher nutrient contents during either the spring or fall. Prairie (27) and alluvial stream bottom and gravel bar (18) settings contain the highest number of edible tuber species. Open woodlands (15), glades (14), and marshes, sloughs and ponds (14) contain moderate numbers of species with edible tubes, while other habitats have significantly lower numbers.

One hundred and six plant species with edible seeds have been identified. Alluvial stream bottoms and gravel bars are the richest habitats (55 species) for edible seed plant species, followed by marshes, sloughs, and ponds (45), and prairie (40) settings. Glades (28), and sandstone/limestone ledge and bluff (26) settings, and disturbed grounds contain moderate numbers of species, while other habitats have lower numbers of species with edible seeds.

Plant species with edible leaves, shoots, flowers, and buds tend to be relatively abundant ($n=286$ species) within the Gulf Coastal Plains. The largest number (46) of species is found in alluvial stream bottom and gravel bar settings, prairies and open fields (44), and open woodlands (42). Glades (34), thickets (27), marshes, sloughs, and ponds (17), sandstone/limestone ledges and bluffs (17), xeric habitats (16), disturbed grounds (15), and barrens (14), contain moderate numbers of resource types. Other habitats (e.g., wooded slopes and slope bases and waste grounds) are home to few species with edible leaves, shoots, flowers, and buds.

Seventy-eight plant species and varieties produce edible fruits are present within the Gulf Coastal Plains at large. The highest numbers of species are found along the borders of open woodlands (19), and in alluvial stream bottoms and gravel bars (18). Prairies (15), sandstone/limestone ledges and ridges (13), glades (12), and the borders of thickets (10) have moderate numbers of fruit-bearing species. Other habitats have few or no edible fruit-bearing species.

Twenty-three nut producing species are present in the broad coastal plains region. Of these, 13 are nut-bearing (e.g., hickory, walnut), and 10 are acorn producing species.

Although these species begin production in the spring, most yields are not ready for harvest until the fall. Rich alluvial river bottom settings (12 spp.) and limestone and sandstone ledges and slopes (12 spp.) are the two most preferred habitats. Eight species are found primarily on wooded slopes and in the more mesic settings at the bases of slopes. Nut-bearing species are less common in other habitats.

Fruit and pecan yield statistics between 1919-1989 (Texas Crop and Livestock Reporting Service, 1984:17 see also Hall 1995b:643) show that (1) the Central Coastal Plains counties produce extremely high quantities, and (2) pecan production is characterized by cyclical highs and lows. Years with high nut yields are followed by 1-3 year periods of lower yields. Annual nut yields also are affected by more proximate factors such as the timing and amount of rainfall, the timing of the first and last frosts, defoliation caused by various agents, and insect infestation and disease (Hall 1995b:641; Rector 1996). An example of the effect of excessive rains could be seen in 1985, when rainfall during November and December hampered harvesting activities and caused the sprouting and loss of some nuts on the ground.

In addition to annual fluctuations in yield, the seasonal availability of nuts to humans is dependent on various factors including the length of the ripening period, the date or period when the nuts fall, and the length of time the nuts are on the ground and are available before they are consumed or are carried away by competing animals (Talay et al. 1984). Observations of dates of nut dropping in hickory and walnuts in Indiana indicated considerable similarity between individual trees (Talay et al. 1984). Texas pecan harvesting activity data for 1984 and 1985 indicate that harvesting tends to begin about the tenth of October and the bulk (about 50 percent) of the crops are harvested in November and early December. These patterns indicate that a considerable crop is usually available relatively soon after the ripening date of a particular species. A large number of animals, such as squirrels, deer, bears, a variety of small rodents, and even turkeys consume nuts. Of these, squirrels are the most significant competitors because they begin consuming nuts while they are still green and on the trees. There is evidence that most squirrels tend to consume thinner shelled, larger nut species first before switching to the consumption of thicker-shelled smaller species (see Talay et al. 1984:347). This type of selection might also occur among smaller mammals even less capable of gnawing through thicker nut shells. An additional potential source of nut loss, particularly among acorns, are the numerous acorn-attacking insects that infest a crop particularly while on the ground. Both the potential losses due to mammals and insects place a premium upon aboriginal efforts to closely monitor nut availability and be in the nut producing patches immediately prior to the nut dropping or soon after.

Selected species distribution maps provided by Little (1971), and species ranges given in the checklist of vascular plants indicate that hickory (*Carya texana*), pecan (*Carya illinoensis*), Blackjack oak (*Q. marilandica*), bur oak (*Q. macrocarpa*), and Coastal live oak (*Q. virginiana*), have the broadest habitat ranges occurring in river bottom settings in the coastal prairies and marshes, oak-hickory forests, and river valleys crossing the prairie. Other species have a more limited habitat range, occurring mainly in one or two vegetation regions of the state (e.g., Mockernut hickory in the oak-hickory forest, Shagbark hickory in the coastal prairies and marshes, and the three species of piñon found only on the Edwards Plateau.

When the seasonal availability of all of these plant resource types is considered together and across the Gulf Coastal Plains, it becomes evident that the highest number of edible species (216) are available in the fall. Spring and summer resources include somewhat lower and nearly equal numbers of resource types (171, and 163, respectively), while the smallest number of edible species are available during the winter (44). These figures support the general impression derived from ethnohistoric records (see Chapter 7) that the leanest season during the annual subsistence cycle was winter and the most productive one was fall.

The overall richness of the habitats, in terms of numbers of resource types, indicates that alluvial stream bottomland and gravel bar settings are the richest (149), followed by prairies (127). Open woodlands (94), glades (90), and marshes, sloughs, and ponds (77), have moderate numbers of species with edible parts. Smaller numbers of edible plant species are found in xeric habitats (60), thickets (56), disturbed grounds (46), wooded slopes (40), and barrens (37). The smallest number of species with edible parts is found in waste grounds. Importantly, while alluvial stream bottoms represent the single richest habitat, all upland settings (e.g., prairies, woodlands, glades, ledges and bluffs) combined have a higher number of edible species than riverine habitats. Importantly, the exploitation of these resources would require greater mobility compared to the exploitation of the single riverine habitat.

Summary of Resource Structure

Based on the brief sketches of resource types and their availability presented earlier, it is now possible to summarize the resource structure that characterizes the Gulf Prairies and Marshes region. A mosaic of small inter-fingering biotic and biological assemblages on the coast offers a high diversity of resources in close proximity to each other. A high diversity of aquatic resources such as fish, molluscs, offshore birds

and their eggs found in the bay-estuary-lagoon habitat and to a lesser degree in the beach-prairie fluvial interface and beach-beach proximate habitats, represent resources that can be procured at lower energy costs. The primary biomass (submerged grasses, algae, etc.) in these settings is less accessible to human consumers but the secondary biomass is rich in density, species diversity (i.e., molluscs, fish) and is highly accessible with low energy expenditures, with the possible exception of birds.

Seasonal changes in ambient temperature, water temperature, salinity and the life cycles of species will impact the resource base in predictable manner by either reducing the density of potential prey within the overall region (i.e., seasonal migrations of shorebirds) or changing the species composition, density and/or age profile of taxa within specific habitats (i.e., estuaries, bays, marshes) and overall within the region (e.g., seasonal migrations). In addition to these seasonally predictable fluctuations, unpredictable weather phenomena such as sudden drops in temperature associated with northern cold fronts or high rainfall events can result in major die-offs of a number of species that have a narrower tolerance range (i.e., molluscs and fish). Such die-offs may have provided very abundant although less predictable resources for populations exploiting the coastal bay-estuary-lagoon habitat.

Long-term and relatively gradual changes in climatic conditions also can be expected and should result in major spatio-temporal reorganizations of biological assemblages within the bay-estuary-lagoon habitat in particular. For instance, the formation of the barrier islands off the coast of the Gulf of Mexico has significantly affected salinity levels within Corpus Christi Bay and its secondary bays such as Oso Bay. We can relatively safely assume that short term predictable phenomena, short term unpredictable episodes, as well as long term gradual changes in climatic condition are likely to have reoccurred repeatedly within the region over time resulting in different scales of adjustments to the changing resource base by hunter-gatherers utilizing the bay-estuary-lagoon habitat.

Riverine corridors, such as those that are found at the center of the prairie-riverine habitat offer linear arrangements of diverse biological assemblages cross-cutting otherwise homogenous terrestrial upland settings. The primary biomass of prairie-riverine habitats is high as a product of nut-bearing tree communities, a diversity of tubers and bulbs, and seed-bearing plant species. With appropriate food processing techniques and technology, this biomass is accessible to human consumers. However, the processing costs involved in making at least some of the biomass (i.e., tubers) consumable is energetically expensive. The secondary biomass of such riverine corridors is relatively high as well (e.g., deer, etc)

but some of its components (e.g., squirrels) may be less accessible to human consumers. In general, the prairie-riverine plant communities and biological assemblages may be less susceptible to minor, short term, perturbations than the coastal aquatic zone because the vegetation is supported by high water tables. However, some of its most productive species (i.e., nut-bearing species) have natural lows and peaks in year-to-year production that add a certain degree of potentially predictable variation in hunter-gatherer resource procurement. Most of the terrestrial species that gravitate to these prairie-riverine habitats have broader tolerance limits compared to those inhabiting bay estuary-lagoon habitats and the co-occurrence of a mix of both aquatic and terrestrial taxa creates a relatively species diverse context buffered to a certain degree from seasonal fluctuations in resources due to the more moisture-abundant setting. For instance, aquatic resources such as freshwater fish and molluscs are affected only by sudden changes in temperature rather than both temperature and salinity conditions. In addition, stream bottoms provide a broader range of micro-habitat variations (distinct depths of water and sediment conditions) that will allow both fish and molluscs to find appropriate habitat conditions with greater ease than within some coastal bays that cover hectares and have a mean depth of less than one meter.

Long-term changes in climatic conditions (e.g., the onset of xeric or mesic conditions) can result in slow and gradual changes in biological assemblages even in these more stable (e.g., prairie-riverine) habitats. For instance, it has been well documented that extreme long-duration xeric conditions tend to result in the replacement of arboreal communities (i.e., woodlands) by grasslands. In contrast, prolonged mesic periods will result in the expansion of woodlands at the expense of grassland prairie communities. It is expected that riverine biological assemblages found in the Gulf Coastal Plains would be similarly affected by these long-term and persistent changes in climatic conditions. Extended mesic periods would lead to the expansion of stream bottom woodlands onto the prairie margin and the formation of oak motts across the open prairie. Both of these phenomena could potentially lead to increased mast production and potentially greater deer population densities. Mesic conditions may also improve forage quality on the open prairie attracting higher densities of large-bodied prey (i.e., bison) as well as prairie adapted antelope. Extended xeric conditions may increase the preferred habitat for antelope but other large and medium-bodied species may decline in response.

Seasonal changes in climatic conditions also impact the biological assemblages found within the prairie habitat. Grasslands represent a relatively homogenous vegetation community but the presence of scattered stands of oaks and oak motts and ponds and depressions provide niches for a range of wildlife and would be critical to deer, one

of the two medium body mammals present on the prairie. Primary and secondary producers in these settings are well adapted to these seasonal shifts in temperature and available moisture. These ebbs and flows result in predictable changes in the amounts and types of resources available to human consumers. Unpredictable phenomena, such as prairie fires can result in dramatic changes in the types and quantities of available resources. At least some of these phenomena can have long-term beneficial effects on some biological communities (e.g., the artificial deflation of woody invaders, introduction of nutrients into the soil.). Long-term processes and forces can have even more dramatic effects on upland biological assemblages given their relative homogeneity and reduced moisture availability. Prairie margin settings will be subject to fluctuations in community boundaries while the species composition of vegetation communities may change (woodlands vs grass vegetation communities) in more homogeneous settings.

Conclusion

Our goal in providing this review is to identify the principal potential resources and define their spatial and temporal availability as well as their potential contributions to the nutritional intake of hunter-gatherer populations. It would be necessary to compile an exhaustive database of the nutritive value of all the edible species available within the region to conduct quantified analyses of the hunter-gatherer behavior at the level of 41NU2 and the region. Unfortunately, after extensive search of the topical and regional literature, it is apparent that the nutritive values per unit of weight as well as fluctuation between age, sex, and season of available resources are unavailable or deeply buried in the literature for the study area. In some cases, information is available for the same species in a different location but the goal of this research module was to determine the specific costs and benefits of hunting and gathering within this study area. Many case studies elsewhere refer to Watt and Merrill (1963) for nutrition information, finding the nearest applicable item listed (Glassow and Wilcoxon 1988, Reidhead 1981), submit samples for nutritional analysis as part of the research, or rely on other published sources, adjusting data to include factors that would effect the nutritional gain of a given resource. Relying on general data such as from Watt and Merrill (1963) disregards specific nutrition information at the species level, age group, season, and habitat of the resource, a sacrifice most researchers find prudent as finding sources reporting nutrition information specific to seasonal or age fluctuations is even more problematic.

In addition to concluding that nutritional data for most Texas and coastal species that we are targeting is sorely lacking or scarce, we are also forced to conclude that the necessary

information related to population dynamics of many of the potential economically useful species is also lacking or is at best incomplete. For instance, even information related to deer densities across the state is inadequate when it comes to seasonal variability in population densities across the state. Data on population densities of smaller, economically less significant species are even sketchier and quantified data on species that are no longer present in the region (i.e., pronghorn antelope and bison) are entirely absent. The same limitations exist with regard to economically useful plant species as well. More importantly, it is necessary to ask whether it is even worth gathering sufficient data on modern resource distributions across the state, particularly in light of the changes such distributions have gone through during historic times.

While we recognize that historic and modern land use practices and developments have dramatically impacted the densities of economically useful species of plants and animals, it is our sentiment that the distribution of plants and animals is determined by long-term evolutionary relationships between soil types, hydrology and vegetation cover and therefore they are

basic to specific organisms and communities of organisms (i.e., relationship of water table to tree species; relationship between mast production and deer population). There is no doubt that in some parts of the state, developments and modern land use have altered the composition of vegetation communities and removed certain animal species. Yet, based on the basic relationship between mast production and deer feeding habits, we can still be reasonably confident in the assumption that deer population densities were relatively higher on the Edwards Plateau than the South Texas Plains and Brush Country even during prehistoric times. Similarly, some of the seasonal changes in the characteristics of the prey species exploited by prehistoric hunter-gatherers respond to seasonal changes in the broader environment (i.e., seasonal variability in available moisture and changes in temperature). As a result, we can be safe in assuming that seasonal changes in prey characteristics (i.e., variation in nutritional composition of plants, changes in the weight and fat reserves of deer), were as much a factor prehistorically as they are today. The point of establishing baselines in resource distribution is not to define exact number of species that may have been present in a particular setting 500 years ago. Instead, the goal is to define the rankings of the habitats based on relative densities of prey species.

Chapter 9: Modeling, Archeological Expectations and the Archeological Record

Steve A. Tomka and Charles Frederick

In Chapter 7 we identified a theoretical perspective, human behavioral ecology, a family of models, optimal foraging, and a specific model, habitat selection that we propose to employ to investigate the foraging behavior of prehistoric hunter-gatherer populations in the project area. In Chapter 8 we described the structure of the natural and biological communities within the Corpus Christi area and the nutritional contributions of selected resources to hunter-gatherer subsistence. We lumped the multitude of niches present in the vicinity of Corpus Christi Bay into five habitats consisting of distinct relatively homogeneous groupings of biological communities. In this chapter, we develop a region-specific set of expectations regarding hunter-gatherer foraging strategies based on the principles of the habitat model and using the selected optimal foraging models described above. In proposing a series of expectations regarding hunter-gatherer land use and subsistence practices, we also define sets of methodological approaches and analytical tools that will be needed to study the data types recovered from 41NU2 and comparative data from regional sites to search for the predicted/expected patterning.

Applying the Habitat Choice Model to Coastal Hunter-Gatherers

In Chapter 2 we have briefly outlined the spatial distribution of the 36 natural environments and biological assemblages present in the Corpus Christi area. We combined these assemblages into five habitats that contain adjoining and similar economically usable resources. We have already mentioned in Chapter 2 that we perceive of a habitat as an aggregation of biological communities that is larger than what would be encompassed within a resource patch, that contains a relatively homogenous group of resources that is the product of long-term evolutionary forces, and that as an aggregate are more similar to each other based on the niche requirements of the resident biological assemblages than to adjacent habitats. We add here that from the perspective of scale, a habitat is larger than the resource patches that would be exploited during a daily foraging round and more than one resident forager group may be present within it at any one time.

As in the prey and patch choice models, the first step in operationalizing the model is to determine the ranking of the habitats. Sutherland (1996:6-7) uses the concept of suitability along which to rank patches. The patch concept is similar to habitats in our definition (see Sutherland

1996:7) and the suitability concept takes into account a broader range of conditions that affect the life-conditions of an organism. Suitability is organism-specific and in our case we are interested in the suitability of the habitat for hunter-gatherer occupation and exploitation. Many aspects of an environment affect its suitability including prey density, intake rate, availability of critical resources such as water and fuel, etc.

To develop a ranking of the five habitats present within the Corpus Christi area and in other bay-associated settings across the Gulf of Mexico, we use the general methodology of rank order determinations for individual prey. As we mentioned before, in energetic terms, post-encounter prey rank is a quantified measure of handling costs compared to Kcal/hr derived from the consumable portion of the prey. Also as we noted before, at the most basic and simplified level, there is a strong relationship between the body size of the prey and its ranking (see Figure 7-1). While handling costs factor into the final determination of ranking, these components rarely shift the ranking of plant species above animal prey or of one animal species above another except in instances of technological adjustments that significantly decrease pursuit or processing cost (i.e., procurement of fish with baskets; making fish-meal rather than individually processing each fish).

Employing these underlying assumptions, we propose to rank each of the five habitats identified in Chapter 2 based on the density of prey in the largest body size class present in each habitat. Unfortunately, even after a concerted effort to compile such information, at this time we do not have quantified data on actual densities of deer, or other mammals within each of the habitats. Instead, we use the concept of potential return rate rather than actual return rate. The potential return rate is estimated based on the assumption that all other things being equal the density of a species within its preferred habitat will always be higher than its density in marginal environments. So that, areas of higher prey density (i.e., deer) will have higher potential return rates than areas of lower prey density. In addition, while the distribution of taxa often extends beyond the boundaries of each habitat, we can reasonably assume that the densities of any given species will be highest in their preferred habitat and lower outside of it (refer to the discussion of deer distribution in Chapter 2). This assumption is important because we do not have density data for specific prey species and therefore have to rely on differences in relative abundance rather than absolute abundance of prey

of different rank (i.e., different size classes). In addition, because the density of some prey species changes seasonally as already mentioned, the ranking of individual habitats also has the potential to change if seasonal changes occur in the density of the dominant mammalian species that determines the habitat's ranking.

Table 9-1 presents the ranking of the five habitats found within the Corpus Christi area. The rankings are derived from and correlated with the body size of the dominant animal taxa present in the habitat.

Table 9-1. Ranking of Habitats by Potential Intake Rate

Habitat	Rank
Prairie-Riverine	1
Prairie	2
Beach-Fluvial Interface	3
Bay-Estuary-Lagoon	4
Beach and Beach Proximate	5

Based on the abundance of deer, the largest size class mammal present year-round in the river bottoms crossing the prairie, we propose to rank the prairie-riverine habitat as having the highest potential return rate. While large size class mammals (i.e., antelope and deer) are present on the prairie proper as well, their densities are not expected to match those of the prairie-riverine habitat for two reasons: (1) the Gulf Coastal Plains and Marshes are on the eastern fringe and outside the recorded range of antelope therefore their densities are not expected to be as high as further west; and (2) deer tolerate and adapt to more open range but the reduced mast resources and cover will maintain their densities below those in the prairie-riverine habitat. For these reasons, we suggest that the prairie habitat ranks second in potential return rate. It is important to note, however, that the prairie habitat is the preferred habitat of bison and therefore, when bison are present in the region on a seasonal basis or during an extended time, the prairie will move up to the number one rank in terms of potential return rate.

The beach-prairie fluvial habitat (i.e., marshes) is home to no large size class mammals and only one medium size class species (i.e., the beaver) is found there in low densities. In contrast, a diverse group of small size class mammalian species are common in this habitat and in combination with high densities of large fish and large birds they contribute to the third rank in terms of potential return rate. The fourth ranked habitat in terms of potential return rates is the bay-estuary-lagoon habitat on the strength of high densities of small and large fish, shorebirds and waterfowl,

molluscs, amphibians and reptiles that represent small prey packages. The beach-beach proximate habitat is the lowest ranked habitat in potential return rates found in the Corpus Christi area. The principal prey species consist of small packages (i.e., amphibians and snakes) and they occur in relatively low densities compared to the neighboring bay-estuary-lagoon habitat.

Because of the seasonal nature of the biological assemblages, no habitat contains the totality of resources needed by human populations at all times (see Chapter 8). Seasonal climatic changes that precipitate the immigration or emigration of species from a habitat can affect the ranking of the habitat since they can result in increases or decreases in the density of high ranked resources. As mentioned above, seasons of bison abundance on the prairie may cause the ranking of the prairie to increase above that of the prairie-riverine habitat. In contrast, high human population densities and hunting pressures within the prairie-riverine habitat may drive out deer into less preferred neighboring habitats. However, immigration or emigration of species from the other three habitats are not expected to affect their ranking because these migratory patterns are temporally/seasonally synchronized so that bird migrations or the spawning migration of fish will occur and impact all three remaining habitats at the same time. In addition, even with the influx of high numbers of large fish and birds, the ranking of these three habitats will not climb to that of the top two habitats unless deer densities in these other two drop catastrophically. While edible plant resources occur in each of the five habitats, their ranking is so low compared to animal packages that they make no significant impact on the overall ranking of the habitat itself.

Region-Specific Modeling and Expectations

The Regional Land-Use and Subsistence Model

In this section, we define a series of expectations regarding hunter-gatherer regional land-use patterns and subsistence strategies through space and time. While the discussion addresses regional-scale patterns, we will also focus on how these patterns may implicate changes in the use and function of 41NU2 over time.

Earlier in this chapter, we proposed that when bison are present only in low numbers or are absent (i.e., latter part of the Early Archaic and Middle Archaic periods) on the Coastal Plains and Marshes the prairie-riverine habitat would be ranked highest in terms of potential intake rates based on high population densities of deer. Based on the habitat choice model, because the prairie-riverine habitat offers the highest

potential intake rate we would expect that the earliest hunter-gatherer populations coming into the region would situate themselves in this most productive prairie- riverine habitat.

Therefore, we would expect that some of the oldest traces of occupation within the region (i.e., residential campsites) would be found within the prairie-riverine habitat. Daily foraging trips would be contained within the river valley and may access resources upstream and downstream as well as at the ecotone between the woodlands and the grassland prairie. Differences in site size may be due to differences in occupation length, which may in turn be conditioned by the productivity of certain resources within the habitat. Similarly, the presence of hyper-productive resources (e.g., pecan groves, oak motts) may facilitate the formation of seasonal population aggregation sites such as those hinted to in the Cabeza de Vaca accounts (see Chapter 7). Aggregation sites may be identified through the presence of features, facilities (e.g., hot rock ovens) and/or a preponderance of tools employed in bulk processing (e.g., grinding stones), temporary storage facilities for processed resources, and a higher richness of tool stone sources present in the unmodified lithic debitage than in residential sites. Of course, the bulk exploitation of certain resources may not require a specialized tool kit (i.e., the exploitation of prickly pear patches described by Cabeza de Vaca) and therefore may leave few technological or archeological indicators.

During the early stages of prairie-riverine habitat occupation, subsistence practices would be focused on the hunting of deer from residential camps during daily foraging trips. Given high return rates and reduced transportation costs, no field processing of the carcass is expected. As, intra-habitat return rates decline and foraging radii increase, field processing may be adopted with low meat-utility elements being discarded in the field rather than returned to residential camps (Bettinger and Malhi 1997). Low ranked animal species may be included in the diet when captured during daily foraging trips. The species of plants and low ranked animals represented in the faunal assemblage of individual prairie-riverine foraging sites will depend on the season of occupation and the post-encounter ranking of the resource or its hyper-productive status, or may be totally unrelated to its ranking based on return rates if obtained by women and children during daily foraging trips.

The habitat selection model predicts that as hunter-gatherer population densities increase friction between groups will begin to interfere with efficient foraging and depletion of resources will cause the widening of the diet breadth. Although, it is expected that initial increases in diet breadth will result in decreased search costs, over time, as more lower

ranked resources are incorporated into the diet, neighboring habitats may provide higher encounter rates for lower ranked prey types if they represent the prey's preferred habitat. Initially, adjustments in resource procurement may simply consist of increases in daily foraging trip lengths and greater residential mobility (e.g., both the length and frequency of moves). Logistical mobility may also be employed to reach more distant but more profitable resources such as deer, pronghorn or bison. Such increase in range of mobility may reach a point where some of the species that are included in the diet will come to represent a large proportion of the optimal diet. If the preferred habitat of these species is outside of the prairie-riverine habitat, human populations may relocate to the new habitat. In addition, when the optimal diet has broadened to this extent, either as a result of interference or resource depletion, it is expected that newly arriving hunter-gatherer groups will begin to settle in the habitat with the second highest potential return rate.

As the diet breadth widens, animal species that rank immediately lower than deer will be pursued by hunters and included in the optimal diet. Deer will continue to be pursued when encountered but because encounter rates are expected to be depressed within the prairie-riverine habitat, it is likely that their representation in faunal assemblages will be biased by field processing in response to greater travel distances. The logistical procurement of deer and pronghorn antelope from the neighboring prairie may further influence the portions of the carcass represented at residential sites. A broad range of plant resources and small-bodied animal species may be included in the diet to alleviate significant day-to-day and seasonal fluctuations in caloric intake.

We would expect the prairie habitat to be the place to draw the next wave of hunter-gatherers coming into the region. However, characteristics of the prairie proper related to its suitability for habitation, such as lack of freshwater and decreased availability of fuel may detract from the intensive residential use of this habitat. However, heavily wooded patches, such as oak motts found near springs or other sources of freshwater, may provide favorable habitation locales. Nonetheless, given its high potential return rate, we expect that the exploitation of high ranked mammalian species through logistically organized mobility will occur within the prairie habitat. As suggested under the discussion of central place foraging, logistical mobility can be a profitable strategy through which to procure high-caloric-yield resources when their density is below the threshold that would encourage nearby settlement. Cabeza de Vaca (1555) describes what appear to be small logistical groups venturing onto the prairie to hunt deer during three-day forays. In addition, however, communal drives that resulted in the impounding of large numbers of deer are also described (see Chapter 7). These

communal drives that relied on fire to concentrate prey and lead to their impoundment, may have been practiced because of low deer population densities in portions of the Coastal Plains under which individual pursuit or ambush hunting may have only lead to low encounter rates. Communal drives, on the other hand, as we mentioned earlier, can result in high return rates especially when all of the prey secured is seen as the product of a single procurement episode.

Even when bison were present on the prairie and the habitat attained the highest rank in terms of potential intake rate, hunter-gatherer groups may have faced the same habitat suitability constraints. It is therefore possible that even under such circumstances, the intensity of habitat use was less than expected or the organization of habitat use was different from that of the prairie-riverine habitat. That is, the exploitation of the prairie habitat may have taken place under logistical organization rather than through residential mobility.

Population packing in the prairie-riverine habitat in combination with the unfavorable living conditions on the prairie proper may have pushed human populations into lower ranked beach-fluvial interface (i.e., the marsh) habitat. The exploitation of the beach-fluvial interface habitat may have most likely occurred from the nearby prairie-riverine habitat because marshy areas are rarely suitable for human habitation. Nonetheless, the presence of small prey species in conjunction with high encounter rates result in moderate potential intake rates in this habitat. We expect that the exploitation of this habitat would increase as the diet breadth increases in the prairie-riverine habitat and as the resources that are incorporated into the wider diet within that habitat approach the return rates for species present in the beach-fluvial interface habitat. Given the low suitability of this habitat for human settlement, we would not expect that under conditions of increased population packing, new immigrant groups or fissioning groups would flock to the beach-fluvial habitat unless regional population densities reach a level where interference or depletion severely lower the potential rate of intake within the two higher ranked habitats. Given the presence of fish and waterfowl it is likely that this habitat would have been exploited by both male and female foragers. There may have been more overlap in prey species procured by men and women from this habitat than in the previous two because the encounter rate of higher ranked prey species is expected to have been lowest of the three habitats discussed so far.

By definition, the diet of groups inhabiting the beach-fluvial interface habitat would include a preponderance of lower ranked prey species. Nonetheless, deer/antelope may be present since hunter-gatherers are expected to pursue these

species when encountered. Low densities of these animals in this habitat may translate to higher transportation costs upon a successful kill leading to field processing prior to transport. Given the seasonal abundance of migratory waterfowl as well as some schooling fish species, we expect some degree of seasonal emphasis on the procurement of fish and waterfowl. The seasonal procurement of partially submerged tuberous/bulbous plant species should serve as a key supplement to a diet consisting primarily of lower-ranked food items. Again, Cabeza de Vaca (see Chapter 7) describes a concerted effort on the part of hunter-gatherer groups to target partially submerged edible plants as well as the effort necessary to process these species for consumption.

While the processing of some plant species appears to have been extremely energy-intensive, the potential to procure large quantities of low return rate resources (i.e., schooling fish; bird eggs) with minimal or low amounts of energy input may have made this habitat a critical resource at least on a seasonal basis when the diet breadth was very broad due to lowered encounter rates with higher ranked species. The use of technologies that would allow bulk procurement (e.g. netting for fowl, baskets for schooling fish, and fishing nets) even when these resources have subsequent processing costs (e.g., sun-drying salted fish) may have still made the strategy profitable if sufficiently large number of prey were encountered and procured. In addition, the availability of large quantities of resources that have no or minimal processing costs would also have added to the attractiveness of the habitat for foragers. In light of recent research in optimal foraging of small bodied prey species (Ugan et al. 2003; Ugan 2005), it is important to recognize that the potential ranking of schooling fish as a highly profitable resource would be dependent on the size of the fish as well as the technological investment in procurement and processing technologies. That is, the procurement of anchovie-sized fish is not energetically profitable if a great degree of effort is invested in the manufacture of nets and post-encounter processing (i.e., grinding salted fish for fish-meal). On the other hand, however, the harvesting of larger fish and reductions of effort invested in procurement (baskets as opposed to nets) and processing (sun-drying rather than grinding) technology has the effect of greatly increasing the energetic ranking of these resources.

The bay-estuary-lagoon habitat is fourth among the five coastal habitats in terms of potential intake rate. The bay-estuary-lagoon habitat proper is not directly habitable, therefore we expect that typically it was utilized while populations lived on the nearby shores. On a seasonal basis intermittent creeks flowed into the estuaries from the nearby uplands, and may have allowed occupation adjacent to the mouths of such creeks. This would have offered a rare source

of freshwater that appears to have been of short supply in this habitat. Exploitation of the small to very small prey species present in this habitat could have occurred on a seasonal basis when potential return rates in neighboring habitats would have been depressed due to low encounter rates with higher ranked resources. We expect foragers to have been present on the shores of these estuaries on a daily basis. The archeological signatures of such low-intensity and intermittent presence along the shores of estuaries may over time combine into a palimpsest of features and occupation debris. The debris may be the product of short term (i.e., day-camp) use and specialized processing stations if the field-processing of molluscs and fish contributed to the lowering of transportation costs and increased maximum travel distance. At points along the shores of estuaries, where intermittent creeks flowed into these bodies of water, the availability of drinking water may have contributed to longer settlement and systematic reuse during the rainy season.

As in the beach-fluvial interface habitat, within the bay-estuary-lagoon habitat the availability of migratory prey species and schooling fish species has the potential to increase encounter rates leading to highly profitable bulk procurement episodes. Rudimentary technology (e.g., baskets) and minimal post-encounter processing (consuming sun-dried fish) can further increase the profitability of these prey species or resources. The availability of low ranked resources that can be bulk-procured may in fact raise the ranking of the habitat particularly on a seasonal basis (but see Ugan 2005). It is therefore possible that this habitat may have been exploited on a seasonal basis even when human populations were concentrated in higher ranked habitats such as the prairie-riverine zone. Short-term exploitation episodes targeting lower ranked resources harvested/procured in bulk would have left little residential refuse and minimal signs of bulk processing if such processing was rather low intensity (i.e., sun-drying).

The beach and beach-proximate habitat is the lowest ranked among the five defined on the coast. It has a low potential return rate but portions of it are in the vicinity of the bay-estuary-lagoon habitat and therefore it may serve as staging grounds for short procurement trips into the estuaries or processing stations for resources identified for transport. For instance, Cabeza de Vaca (1555) indicates that groups inhabited the barrier islands and the account suggests that access to some parts of the coast may have been limited to certain groups. We do not know whether this reflects the presence of territorial groups on the coast but beachfronts, more than many other resource configurations, may be territorially defendable at reasonable energetic costs. Groups along beachfronts may not have owned the resources within these bodies of water, but control of the access points to those resources may have

at least limited where the resources were harvested from. The large prickly pear patches and highly productive, though cyclical, pecan forests also were not owned as far as we can establish from the accounts reviewed. Regardless of whether access was controlled or not, the presence of cemeteries on the coast as well as inland in the vicinity or within pecan forests may be related to the development of territoriality and may be intended to mark territorial boundaries or possession of a spot on the landscape. Alternatively, rather than reflecting territoriality, the presence of cemeteries may be a signal of usufruct rights to specific parts of the landscape, habitats and perhaps patches of resource.

Hypothesized Changes in Function and Use at 41NU2

Having presented a broad model of regional land use and subsistence strategies by coastal hunter-gatherers, we can now present a series of hypotheses regarding the role the Cayo del Oso site may have played in this model, how the role may have changed over time, and how the material remains found at the site may be reflective of the pattern of site use and subsistence practices at 41NU2, the Cayo del Oso site.

In Chapter 1, we described that the site sits on the shore of a shallow estuary bay, False Oso Bay, less than one kilometer from Corpus Christi Bay. A small intermittent creek, the False Oso Creek drains into False Oso Bay near the southern end of the site. The location fronts the Bay-Estuary-Lagoon habitat and the Prairie habitat surrounds it on land. These habitat conditions have most likely changes dramatically over time significantly influencing the resource base and structure in the vicinity of the site. Ricklis focuses on the development of estuarine habitats as the primary factor conditioning an intensified use of the margins of Corpus Christi Bay after about 3000 years B.P. Nuclei of transgressive barriers may have begun establishing estuarine conditions within the bays of the coast even as early as 6800 B.P. although changes in sea level may have influences the productivity of proto-bays and lagoons favoring human exploitation or abandonment as suggested by the hiatus of C-14 dates between 6800-5900 and 4200-3000 B.P. In addition, Albert's (2006) analysis of pollen spectra from Swan Lake and Nueces Bay indicates that several periods of sub-tropical to tropical conditions may have alternated with more mesic conditions between 7100 B.P. and modern conditions leading to shifts between rather wooded and more open grassland terrestrial conditions in the Copano and Nueces Bay areas. We mentioned earlier that the prairie habitat may shift in potential return rate between the top two ranks dependent on the presence or absence of bison. We also indicated that certain features of the prairie habitat, such as lack of drinkable water, lack of firewood, make it less attractive for human habitation unless oak motts

and sources of drinking water can be located. Finally, we also mentioned that mast production may be one of the key resources influencing deer population densities. Given the effect of these changes to both the estuarine habitat and to the structure of the resources and the potential return rate of terrestrial habitats, it is critical that the climatic conditions and vegetation communities present in the vicinity of 41NU2 be defined. Depending on which climatic/vegetation conditions prevailed and how they shifted over time, several potential reconstructions of the occupation history of 41NU2 are likely.

If estuarine conditions were not fully developed until sometime after 3000 B.P., prior to this date human populations would have likely concentrated in the prairie-riverine habitat located well upstream of the site. Proto-estuarine conditions may have resulted in increased biomass that may have been accessible to human populations from the shores of Corpus Christi Bay using specialized technologies such as canoes and netting. Therefore, under these conditions we hypothesize that human occupation of 41NU2 would have occurred primarily during periods of higher effective moisture and would have been of a short-term extractive nature where foraging groups would have reached the locality through overnight foraging forays from inland residential camps located in prairie-riverine settings. The distance from the mouth of the Nueces River to the site is approximately 35 km, slightly outside of the typical daily foraging distance. During such extractive forays, gathering groups would have concentrated on the procurement of low ranked resources such as birds, fish and molluscs.

From this hypothesis we can derive a series of specific expectations regarding the nature of change in occupation of 41NU2. Specifically they include the following:

- (1) the earliest occupation of 41NU2 may well predate the formation of the barrier islands therefore the oldest dates should fall prior to 3000 B.P.;
- (2) the material remains associated with these early occupations would consist of limited debris from the extraction and processing of small quantities of resources at day-camps (i.e., shell concentrations; small thermal features);
- (3) resource extraction would have focused on fish, mollusks and birds;
- (4) in the absence of suitable watercraft and technology to allow the efficient exploitation of deeper waters,

molluscs and fish otoliths would reflect the exploitation of shallows as opposed to deep aquatic environments;

- (5) repeated seasonal use of the site would lead to palimpsest occupation surfaces resulting from multi-year reuses of the locality;
- (6) extended periods of higher effective moisture on a seasonal basis would lead to systematic re-occupation of the site, the accumulation of abundant organic debris from site use, food preparation and discard; and
- (7) such periods should correspond to higher richness in prey species representing the systematic exploitation of both terrestrial and aquatic species and greater density of features;
- (8) an increase in deer skeletal specimens in the faunal assemblage should correlate with such occupations;
- (9) the increased role of hunting would result in higher numbers of projectile points in assemblages;
- (10) seasonal short-term occupations of the site should generate lithic tool assemblages characterized by low richness due to the relatively narrow range of activities carried out on site; and
- (11) lithic resources would be dominated by materials obtained further upstream where stone tool materials are more common and of larger size.

We expect that until the establishment of the hyper-productive bay-estuary-lagoon system or major shifts in regional climatic regimes, the exploitation and use of 41NU2 would be characterized by short-term intermittent (i.e., opportunistic) and systematic seasonally recurrent strategies. The focus of residentially mobile populations would be the prairie-riverine habitat and the exploitation of the 41NU2 site would be in a supportive and seasonal role. We suggest that the preferred seasons of visits to and exploitation of the coast (i.e. the area of 41NU2) would potentially be conditioned by the resources being targeted for extraction. For example, fall and spring exploitation of the coastal belt would provide access and opportunities for the en-mass extraction of schooling fish and congregating birds. Exploitation of coastal areas near the ends of the seasonal rainy seasons would increase access to fish populations that may be trapped in shallow pools of water as these dry up and become less numerous. Winter and mid-summer (i.e. around July) exploitation of the coastal

habitats would provide easier access to molluscs and the few winter-spawning fish that would be found in the shallows.

Over time, increased human population densities in the prairie-riverine habitat would result in lower potential intake rates through interference and depletion of resources. How long it would take for a noticeable decrease in resources depends on the density of human population and the density of resources within habitats, as well as a variety of other factors including extractive technologies employed, mobility strategies practiced and the potential of resources to regenerate. The depletion of resources will favor the “colonization” or movement of new or fissioning population segments from the prairie-riverine habitat into the vicinity of the bay-estuary-lagoon habitat and to the area of 41NU2. Although such a population influx into the habitat may happen even before the establishment of full-blown estuarine conditions (i.e., the formation of barrier islands), it is most likely that the bay-estuarine-lagoon habitat will reach its highest potential return rates once the estuaries are established and high effective moisture conditions prevail across the region. We perceive this full-scale shift of human populations into the habitat as triggered by population packing and reduced potential return rates in higher ranked habitats and the increase in potential return rates in the bay-estuary-lagoon habitat brought on by the establishment of estuarine conditions and long-term high effective moisture regimes (i.e., mesic climatic regime). The change in land use orientation would be brought on by resource depletion in the prairie-riverine habitat and reinforced by the increase in potential return rates in the bay-estuary-lagoon habitat. Such a shift would represent a dramatically different use of the habitat and the site itself. That is, in contrast to the previous pattern, the bay-estuary-lagoon habitat and its proximity will represent the focus of year-round resource extraction and mobility with logistical moves outside of it to obtain higher-ranked resources from neighboring habitats.

We propose the following series of specific hypotheses regarding this shift in the nature of occupation of the bay-estuary-lagoon habitat and of 41NU2, itself:

- (1) the shift in intensity and the nature of occupation of 41NU2 should occur as a result of out-migration of hunter-gatherer populations from the prairie-riverine habitat and would be made feasible by the formation of the barrier island and therefore should post-date 3000 B.P.;
- (2) the archeological remains associated with these occupations should consist of abundant debris derived from a combination of extraction and processing of

seasonally available hyper-productive resources (i.e., molluscs, fish and birds) and residential occupations associated with seasonal exploitation of the resources found in the habitat;

- (3) resource extraction would have focused on fish, mollusks, birds and small to very small terrestrial species secured during daily foraging trips;
- (4) some hyper-productive resources found within the habitat (i.e., schooling fish) may provide systematic opportunities for bulk extraction and this in turn may lead to the construction of facilities (i.e., drying facilities) and use of technologies (i.e., grinding) that facilitate bulk processing (i.e., fish meal) and perhaps even storage facilities (see Ugan 2005, however);
- (5) highly ranked resources such as deer, antelope and bison would be obtained through logistical excursions into the prairie-riverine and prairie habitats;
- (6) due to the greater transportation costs involved, skeletal elements of highly ranked mammalian species found at 41NU2 would be biased in favor of high-meat utility elements as lower ranked elements are discarded during field processing;
- (7) as a result of longer seasonal occupations, a greater variety of activities would be carried out on-site leading to increased richness in feature types;
- (8) similarly, an increased richness in functional tool forms is expected as a result of the greater variety of activities and longer site occupation;
- (9) as in the previous instance, lithic resources would be dominated by materials obtained further upstream where stone tool materials are more common and of larger size; the procurement of such resources would be imbedded in the long-distance logistical forays;
- (10) because of longer term use of the habitat, however, it is expected that the stone tools discarded under such occupations would be more heavily reduced and a greater emphasis would be placed on rejuvenation and recycling of artifacts.

Albert's (2006:172) pollen spectra suggest that the modern still stand period from approximately 3050 B.P. to modern

times is characterized by high effective moisture levels. The formation of full-scale estuarine conditions sometime after 3000 B.P. would have created conditions that favored intensive utilization of the bay-estuary-lagoon habitat, particularly if population packing in the prairie-riverine habitat had a negative effect on potential return rates for hunter-gathers. As a result, the depositional zones accumulated following 3000 B.P. should be rich in organic materials, artifacts and features and such zones should not oscillate with less organic depositional zones because the favorable resource conditions would have persisted during the remainder of the Late Archaic and Late Prehistoric periods, according to Albert (2006).

Finally, population packing, high population densities, and the potential to control access to seasonally hyper-productive resources may have brought together a series of factors that favored the establishment of some degree of territoriality and the control of access to coastal resources by limiting access to stretches of the coast. If such was the case, the establishment of cemeteries may have potentially served as a strong symbolic marker of a certain usufruct-right to the resources of a locality, if not out-and-out ownership of the access routes to those resources. On the other hand, the multiple sites containing cemeteries within the Coastal Plains region may simply be the product of repeated revisiting and extended stays in localities endowed by a predictable food supply.

Defining Research Problems, Data Types and Analytical Approaches

Unlike typical research designs where much effort is spent in identifying analytical units, in our case, given the nature of the monitoring project, we will work with features and aggregates of features (i.e., multiple features composed of distinct, thin scatter of shell, small flecks of charcoal and burned clay, and animal bone, all confined to a specific paleosol) that form occupation zones, as the basic analytical unit. Therefore, the discussion of research problems and analytical approaches that follows is not tied to the typical analytical units expected to derive from most data recovery projects.

The discussion of the research problems begins with a brief statement regarding the relationship of the needed research to the overall study. This is followed by a number of specific research questions for which answers will be sought and the listing of the analytical approaches that were proposed in the research design to answer the questions and address the overall research problem. Each section is closed with a brief discussion of the data types used and the actual analyses performed.

Research Problem: *The Need For a Site-Specific Paleoclimate Record*

Earlier in this chapter, we presented a general model of the land use and subsistence practices that we expect to have occurred on the middle portion of the Gulf Coastal Plains and Marshes. We also developed a series of specific expectation regarding changes in site function, site use and concomitant material correlates that we should encounter in the materials recovered from 41NU2. Underlying these specific expectations are three key factors, oscillations between tropical and moisture rich periods, human population packing and concomitant reductions in return rates within given habitats, and the formation of barrier islands, that have influenced and conditioned the types of prehistoric resources and their structure which in turn influenced the types and nature of human occupation of 41NU2.

Given the critical role of environmental conditions, and because Albert's (2006) analysis is the first and only region-specific reconstruction of terrestrial conditions, we propose that one of the first research needs is the development of a site-specific paleoclimatic sequence.

Associated Research Questions:

- (1) What is the paleoclimatic record as manifested at 41NU2 and how does it compare with sequence established by Albert (2006) and does it have any implications for the periodic sea level changes suggested by Ricklis (2004)?
- (2) Do the trends in site-specific paleoclimatic conditions correspond with changes in the structure of the economically usable resources proposed in the broader regional model?
- (3) Do the oscillations in sparse- and heavy-organic content strata at 41NU2 reflect periods of rapid and slow deposition, respectively or do they represent different intensities of site use?

Proposed Research: A variety of analytical approaches and archeological data types can provide primary or secondary proxy information on climatic conditions. Those approaches and data types that we hoped to recover from 41NU2 are briefly reviewed below.

Pollen, primarily from bogs and dry caves in Central, North, and far West Texas, has been used extensively to reconstruct past climates within the state (e.g., Bousman 1994, 1998; Bryant and Shafer 1977; Camper 1991; Potzger and Tharp

1943, 1947). Under conditions of constant moisture, slightly acidic soils, and rapid burial, pollen preservation is often excellent, and when dates are available, pollen can provide detailed reconstructions of vegetation communities. Unfortunately, pollen does not seem to preserve well in open sites, in locations with alkaline sediments, or locations that experience frequent shifts in moisture. In these settings, pollen samples are likely to be characterized by low grain counts and selected preservation of certain, hardy spores (see Bryant and Holloway 1985; Bryant et al. 1994; Holloway 1989). While characterized by alkaline sediments, the overall conditions at 41NU2 suggest that pollen may be preserved, especially in the pond sediment. Although preservation of pollen within the dune may be less likely (as Kibler 1994 found), pollen has been found preserved within clay dune deposits elsewhere (e.g., Rhodes and Rich 2004) and it would be worth checking the buried soils within the dunes to see if pollen has been preserved. If this is the case, the analysis of pollen from this region may prove to be a significant addition to our understanding of past climate in this region.

Diatoms are microscopic, unicellular photosynthetic algae that are found in a wide variety of damp terrestrial and aquatic settings. With silica cell walls, diatoms are resistant to many chemical and mechanical processes that differentially destroy pollen. As such, they have high recovery potential in many archeological settings. As a group, diatoms appear to be found over a wide range of environments (see Battarbee et al. 2001; Patrick 1977; Owens et al. 2004). While not extensively used in Texas, there are several studies within the state that attempt to use diatoms for the reconstruction of past environments (e.g., Winsborough, 2004, 2007). Diatoms are used extensively in paleoenvironmental reconstruction in other regions (e.g., Weckstrom et al. 2006). Several individual species of diatoms have specific tolerances to factors such as pH, salinity, pollutants, water temperatures, and water depth, and the presence of specific species, as well as shifts in the number and overall richness of species (see Bloom et al. 2003; Cumming and Smol 1993; Kindt and Small 2002). As such, an analysis of diatoms may provide important data on local as well as regional conditions. There is a high probability that diatoms will be recovered from the cores, especially in the pond setting. If so, their analysis may potentially provide direct information on moisture regimes, temperature, and salinity levels, all of which have implications for resource structure.

Opal phytoliths are, in effect, copies of plant cells and intercellular spaces that are produced as a result of plants ingesting water that has dissolved silica present. Phytoliths form in a variety of plants, though they are best studied in grasses. A wide variety of shapes may be associated with a given species, and roughly similar shapes can occur in

unrelated species. Nevertheless, it appears that classes of shapes and relative frequencies of phytoliths can be used, in several cases, to identify grass and other plants at the species level (Ollendorf et al. 1988; Piperno 1988; Rovner 1983). Several phytolith studies focused on reconstructing paleovegetation, and by extension paleoclimate, have been undertaken in Texas. Robinson's work (1979, 1982) in coastal and near coastal settings has been discussed previously. Additional studies have been conducted in the Rolling Plains region of Texas (Fredlund et al. 1998), central Texas (Fredlund 1994), and in far west Texas (Buck 1993). Phytoliths share some similarities with diatoms in that they are composed of silica and are, therefore, resistant to decomposition. As such, they have a high probability of being recovered from most contexts, and there is a high probability that phytoliths will be recovered from our cores. If so, their analysis should provide direct information on vegetation present, as well as changes in that vegetation. In that regard, phytolith analysis would prove complementary to the analysis of pollen by providing a direct measure of local vegetation shifts.

Ostracods are small aquatic crustaceans that secrete a bivalved shell made of low magnesium calcite and they are a common component in aquatic deposits in general (cf. Holmes and Chivas 2002a; Boomer 2002) and have been previously examined in this region (cf. Cronin 1986). Although the paleoenvironmental record provided by ostracods is often less detailed than that offered by diatoms, these microfossils are quite easily used to distinguish between major environmental shifts such as are locally present (e.g. marine, lagoonal and freshwater) and these microfossils may be preserved in the pond sediment. Although the late Holocene changes in ostracod fauna within the pond are expected to be minimal, there may be a climate signal associated with drier periods, when more marine taxa may be present owing to eolian transportation. For this reason, it may be worth characterizing the pond ostracod fauna. A second reason to examine the ostracods is that their shells may be used to provide a chronology. The biogenic calcite shell or test of ostracods has been successfully radiocarbon dated previously and appears to yield reliable radiocarbon ages (Holmes and Chivas 2002b:199).

In addition to microfossil proxy records of environmental and climatic change, the lithological sequence of dune growth and stability, as denoted by the presence of alternating bands of light colored pedogenically unaltered dune sediment, and weakly developed buried soils within the clay dunes may provide a record that is climatically sensitive. Clay dunes, lunettes, or lomas, as they are locally known, are a form of source-border dune that are formed from the deflation of sand sized aggregates of mud from fringing mud flats. The production of sand-sized clay aggregates that are available

for wind transport is facilitated by salt crystal growth in the drying mud flat. The driving force in clay dune formation is clearly one of local hydrology (cf. Bowler 1973; 1983) but studies of some dunes (e.g. Holliday 1997:67) have concluded that because local hydrology responds to region climate, the lithological record of such dunes may have regional paleoenvironmental significance. Huffman and Price (1949:120) note that the formation of clay dunes on the Texas coast is strongly related to climate:

“In south Texas, movement of the clay pellets begins in the month of March and continues until November when the growth of the dunes is arrested by winter precipitation and reduced evaporation. Their growth is most rapid during the drier summer months and during periods of prolonged drought.”

An example of how much sediment may accumulate in a drought was provided by Price and Kornicker (1961:249) who noted that in the 7 year drought of the 1950s, some dunes accumulated a foot or more of fresh sediment. Price and Kornicker (1961:246) have also noted a climatic influence on the size of clay dunes with the dunes being highest near the Rio Grande Delta and lower toward the progressively more humid climates to the north and south. Given these observations it stands to reason that the periods of dune growth and development may provide a climatically sensitive record, whereby the dune grows during periods of relative aridity and become stabilized during more moist periods when soils form. Erosion of dune surfaces may also occur during periods of high effective moisture or isolated rain events that occur during dry periods as suggested by erosional features detected across the northern and southern dunes at 41NU2.

One way of testing this hypothesis is to compare the stratigraphy of the two different dunes, and radiocarbon date the buried soil A horizons. Kibler (2006, personal communication; 1994) obtained a sequence of radiocarbon dates on bulk soil A-horizons within clay dunes near the Rio Grande delta and the dates appeared to yield a reasonably good chronology that was consistent with the interbedded archeology. If climate is the driving force behind periods of dune aggradation, then the two dunes can be expected to yield comparable paleosol sequences (in terms of the number of buried soils and their radiocarbon age). If, however, there are significant differences in the soil stratigraphy between the dunes, then other, possibly local, factors other than climate may be critical in the sequence of growth and stability of these clay dunes. Furthermore, if the sequence of dune growth and stability is a regional climate record, the record

from the Oso Bay dunes should bear a strong similarity to the record obtained by Kibler (1994) in Hidalgo and Willacy County, near the Rio Grande Delta. Finally, it was hoped that if a pollen record can be derived from the pond, comparison of the dune lithological record with the pollen record may provide a useful test of the hypothesis.

Because the primary goal of the 41NU2 monitoring project was to rapidly document features and exhume human remains when encountered, none of the data types listed above were available to CAR at the end of the monitoring project to adequately address the reconstruction of the paleoenvironment at 41NU2. Therefore, in the subsequent research design, we proposed to recover six soil columns, two each from the northern and southern dunes on site and two from the buried pond deposit immediately northwest of the site.

At the time of the writing of the research design we also proposed to conduct macrobotanical analysis to identify elements of the vegetation community that was present in the vicinity of 41NU2. In addition, we hoped that the macrobotanical analysis would identify charred mesquite fragments. Dering (2002:206-7) has developed a technique for measuring the amount of rainfall available during the life of mesquite trees, by studying the variations in the anatomical characteristics of the xylem preserved by charring. Noting that mesquite can adapt to ranges of annual rainfall average of from 20 cm to 102 cm, Dering (2002:207) states “such adaptability requires a certain plasticity of growth form, and it is these adaptive responses to precipitation that we wish to measure”.

Finally, to compliment these data types, we also proposed to carry out the systematic analysis of shell species composition of feature fill from distinct occupation zones as well as of fish otoliths recovered from floated feature fill samples to obtain a profile of changes in species over time. We hoped that such trends would reflect changes in nearby habitat conditions for those organisms. While we recognize that the mollusc and fish species present in thermal features have been influenced by the filter of human selection, we hoped that strong consistent patterns in mollusc and/or fish species representation also may be reflective of changes in habitat conditions (i.e., salinity, stratum).

In proposing these wide ranging research approaches focused on paleoclimatic reconstruction, our hope was that the analyses of multiple independent data sets would provide a more robust foundation from which to reconstruct paleoclimatic conditions. The different data types have different paces and time-lags in responding to changes in climatic conditions and

as such their composite may provide a more accurate picture of when and how climatic conditions changed on the coast.

Available Data and Analyses Conducted: To systematically investigate changes in paleoclimate and to establish a site-specific record, TxDOT-ENV accepted CAR's proposal to return to the site and extract the sediment cores that would be critical to the investigation of the paleoenvironmental research problem.

In August 2007, using a Eijkelkamp percussion drilling set and its accompanying gouges, the project geomorphologist, assisted by CAR staff, extracted three soil sediment cores, one each from the pond, and the north and south dunes. Subsequently, the geomorphologist subsampled the cores/columns to produce samples intended for the analysis of diatoms, pollen, phytoliths, ostracods and a variety of geomorphological aspects. Humate samples were also pulled to allow the dating of the depositional sequences. The potential contribution of each of these specialized analyses to paleoenvironmental reconstructions is briefly outlined below.

To obtain secondary proxy data on paleoclimate heavy fraction from 66 flotation samples of feature fill were submitted to Dr. Phil Dering of Shumla Archeobotanical Laboratory for macrobotanical analysis. Unfortunately, while we had recovered a large number of charcoal samples during the project, and the macrobotanical analysis did identify some mesquite pieces, much of the charcoal consisted of small fragments that were too small to allow for the anticipated xylem analysis. Some of the larger pieces of charcoal that were recovered from the site were not feature-associated and therefore we chose not to submit them for xylem analysis.

The final dataset used in this analysis consisted of the molluscs and otoliths that were extracted from the same 66 feature fill samples floated from the macrobotanical analysis and five additional proveniences derived from occupation zones. In total, therefore, the mollusk and otolith samples from 71 heavy fraction samples were used as proxy data for paleoclimatic interpretations.

Research Problem: *Dating of the Human Occupation Sequences Present At 41NU2*

In addition to constructing a regional paleoclimatic sequence, we also need to be able to anchor the climatic phenomenon and human occupation sequences at 41NU2 in time. The earlier presented model and site-specific expectations have made it clear that we expect significant changes in hunter-gatherer habitat preference and site use

over time in part due to changes in potential return rates and broad population dynamics and their influence on habitat ranking. It is clear that dating the occupation sequences present at 41NU2 is critical in terms of the general model and the site-specific expectations.

Associated Research Questions:

- (1) What are the ages of the components identified at 41NU2?
- (2) Are there components that predate the formation of the coastal barrier islands indicating that human occupation began prior to dramatic changes in estuary conditions?
- (3) How much time elapsed between the formation of organic-sparse and organic-rich zones at 41NU2?
- (4) Are dates obtained on oyster shell reliable in the absence of charcoal dates?

Research Proposal: As a primary goal of this research problem, in the original research design, we proposed to submit feature-associated charcoal samples from the tops and bases of each of the organically enriched strata identified at 41NU2. We also proposed to date multiple features from each occupation zone to ascertain the formation history of the zone. Finally, we hoped to be able to submit multiple samples from a select number of features to determine how long the features may have been in use.

This initially proposed sampling strategy resulted in the identification of 110 samples to be submitted for radiocarbon dating. It quickly became evident that it was financially prohibitive to submit this number of samples. Therefore, upon meeting with TxDOT-ENV staff and discussing the goals and scope of the dating strategy, CAR in collaboration with TxDOT-ENV revised the dating approach.

Available Data and Analyses Conducted: The final dating strategy consisted of the selection of samples from two sequences of features one from each of the north and south dunes. The features were selected in such a manner that they provided a sequence of dates from the top to the bottom of the stratigraphic column of each of the respective dunes. Using this approach, 40 samples were submitted from features. The samples consisted of 20 pairs, one consisting of charcoal and the other of shell (preferably of the same species). In addition, six other samples also were submitted from occupation zones identified during monitoring to define the ages of these proveniences. Finally, 12 humate samples

were also submitted, four from each of the three geomorphic cores extracted from 41NU2. Two of the cores came from the respective dunes and one was extracted from the pond.

The charcoal and shell samples submitted were considered as paired when the feature-associated charcoal sample derived from soil matrix captured within the valve of a shell or when the charcoal adhered directly to a valve. Such paired samples provide the impression that they are more securely associated with each other than samples derived from within 5 cm of each other from within the same hearth. Unfortunately, we cannot think of any recovery context that can provide the full-proof evidence of direct association implied for instance by a projectile point stuck into a piece of bone, which is what would be necessary to investigate the reliability of dates derived from shell versus charcoal. Nonetheless, the above described criteria were the closest approximation of such parameters proposed by TxDOT-ENV.

Overall, this revised strategy resulted in the submission of 58 samples and allowed for the coarse-grained dating of the stratigraphic sequence noted in each of the two dunes and also contributed to the investigation of the correspondence between charcoal and shell dates (see Ricklis 2004:157-161). In addition, it allowed us to relate the stratigraphic sequences noted in the geomorphic cores to the ages of the occupations within each dune.

Research Problem: *Developing a Quantified Set of Optimal Foraging Models and Expectations*

The model and site-specific expectations are based on the size-related ranking of broad prey/resource classes. For a model and related expectations to be truly quantified and testable, a great deal of information needs to be gathered regarding the caloric yield and density of the prey, costs incurred in harvesting and processing the prey/resource for consumption, and other factors. Such detailed data for the regional resources that would have been available to inhabitants of 41NU2 could not be located at the time of the writing of the research design. However, recent HBE and Optimal Foraging Studies have shown a great deal of promise in the quantified investigation of hunter-gatherers as well as food producers (see Kennett and Winterhalder 2006 and references therein).

Associated Research Questions:

- (1) What are the caloric contents of distinct potentially economically useful species of animal and plant resources present within the Corpus Christi area?
- (2) How does the caloric content of these potentially economically useful species of animal and plant resources vary over the year?
- (3) What is the density of the potentially economically useful species of animals within the distinct habitats of the Corpus Christi area?
- (4) How do the densities of potentially economically useful species of animals vary during the year?

Proposed Research: The development of solid expectations that are regionally specific has to come from actual quantifiable data on prey densities, caloric content of different prey sources, and time-motion experiments that replicate procurement and processing of regional prey using traditional technologies. As a first step toward the development of a database that will allow the quantification of basic prey characteristics, we proposed to conduct a systematic search to compile an exhaustive database containing information on caloric content of the principal mammalian and plant resources that would have been consumed by prehistoric peoples living at 41NU2. During the research design phase of the project, we had established that some of this data already existed and has been gathered by nutritionists for a limited number of plants and animals. We also knew that in other instances, resources would have to be acquired and submitted for nutritional analysis. TxDOT-ENV staff assessed this portion to be too ambitious and suggested that instead of an exhaustive search, we initiate a pilot project to systematically review and compile the nutritional information that already exists to establish what data exists and what are the gaps in information.

Available Data Needs and Analyses Conducted: As part of the revised approach, we set out to compile information on four distinct characteristics of regional food resources: (a) the caloric content; (b) fat content; (c) protein content; and (d) carbohydrate content. Because the faunal resources are ranked so much higher than most plants, and because our habitat rankings are based on the rankings of animals, we focused our search on the nutritional characteristics of prey animals. We collected information on the (a) caloric; (b) protein, (c) fat, and (d) carbohydrate content of animal flesh of all species that may have been employed as a food source prehistorically. Because of the scarcity of data for species inhabiting the study region specifically, often we had to rely on information for the same or similar species from other regions.

Research Problem: Reconstruction of Regional Land Use Patterns

We have discussed earlier that the timing and nature of the occupations of 41NU2 may have been conditioned by various regional-scale dynamics including human population densities, regional climatic conditions, and habitat productivity/profitability. When 41NU2 began to be occupied and how that occupation may have changed over time can be understood only within broader regional context of hunter-gatherer exploitation of coastal resources.

Associated Research Questions:

- (1) Does the prairie-riverine habitat have the oldest archeological components?
- (2) What is the density of archeological sites in the five different habitats?
- (3) What was the periodicity of site occupation?
- (4) How did this site articulate with other sites within the four other habitats in this region?
- (5) Does the repeated occupation and abandonment of the 42NU2 site represent a response to oscillations in regional climatic episodes?

Proposed Research: In the research design we proposed to conduct two scales of analysis, one site-specific and the other regional in scale. At the site-specific level, we propose to examine the available faunal assemblages to compare the species represented with the ones that are expected based on the model presented earlier in this chapter. We hoped that this inventory would provide us an approximation of the range of habitats accessed by hunter-gatherers from 41NU2. Also, we anticipated that the faunal assemblage may help identify the procurement seasons for the prey represented and this should serve as a proxy for the season of site occupation. We also proposed to determine what was the range of raw materials represented in the collection of lithic artifacts from the site. Minimum Analytic Nodules were to be defined through the study of the color and inclusions present in the lithic artifacts.

At the regional level, we proposed to compile information of site distributions within Refugio, Aransas, Nueces, and San Patricio Counties to serve as a regional database for the investigation of land use. In addition, we proposed to collect detailed information on a number of well reported

archeological sites and/or components from within the Coastal Plains and Marshes region.

Available Data and Analyses Conducted: Faunal remains identified in the 71 heavy fraction samples, and all other tools collected during the monitoring project constituted the samples of site-specific materials used in this analysis. Only 33 lithic artifacts were present in the 41NU2 assemblage. The majority are lithic debitage (n=23), and only two are projectile points. In addition, a handful of bone (n=8) and a moderate number of modified shell (n=83) tools also were recovered during the project. These tools, lithic artifacts and faunal remains were the focus of the site-specific analysis.

The small sample of debitage was not grouped into homogeneous categories based on color, texture as was proposed in the original research design. The specimens were for the most part so small that it was felt that they were not necessarily representative of the larger cobbles they derived from. Instead, a number of attributes, including the maximum dimension and cortex availability of each piece was noted to define what stage of reduction the specimen represents. The dart points were examined (i.e., various metric attributes) to document their heavy rejuvenation.

As part of the regional-level analysis, we compiled data on the following aspects of sites present in the four counties: (a) the number of sites found within each county; (b) site size; (c) age of component(s); (d) assemblage content by major material culture categories (e.g., chipped lithics; if available, number and types of features and facilities; faunal remains). Using these data on a GIS platform we linked the distribution of sites to precise locations on maps of the region to quantify the following characteristics of site location: (a) distance to nearest freshwater stream; (2) habitat type where site is located; and (3) distances to the nearest habitats.

Finally, we collected information from a small number of previously excavated and reported archeological sites and/or components within the region to serve as the basis of a comparative analysis of the artifact and faunal assemblages. The 18 sites and/or components were: 41NU11, 41NU184, 41NU221, 41NU266, 41NU267, 41NU268, 41NU276, 41SP15, 41SP43, 41SP120, 41SP136, 41SP148, 41SP149, 41SP153, 41SP156, 41SP177, 41AS3 and 41AS16. Information on these sites and their components is presented in series of technical reports mentioned in Ricklis (2004). The data that was collected included: (a) lithic artifacts present; (b) number and types of worked shell; (c) number and types of worked bone; (d) number and types of ceramics; and (e) mammals and vertebrate and invertebrate taxa present.

The comparative study relied purely on already published information.

Research Problem: *Reconstruction of Subsistence Practices*

The model presented in the first part of this chapter identified a series of hypotheses regarding changes in the manner in which the bay-estuary-lagoon habitat was exploited over time. These changes would have impacted the diet of the hunter-gatherer population as well as the material signatures left behind at 41NU2. Once the effect of differential preservation is allowed for, the faunal and macrobotanical assemblages should play an important role in reconstructing hunter-gatherer diet and resource base at 41NU2.

Associated Research Questions:

- (1) Do the economically useful resources recovered from 41NU2 reflect the diet breadth expected based on the optimal foraging model and expectations?
- (2) What was the role of marine resource in the diet of the prehistoric inhabitants and did this change over time?
- (3) Was there a major change in the diet of coastal hunter-gatherers following the formation of the barrier islands?
- (4) Are deer present in faunal assemblages from coastal sites and, if so, what is the meat utility of the skeletal elements present?
- (5) Are there economically useful plants that are present at the site?
- (6) How does the composition of faunal assemblages vary across the research area as exemplified by comparative assemblages from different habitats?

Proposed Research: As in the case of the land use research problem, in the research design we proposed to conduct two scales of analysis related to hunter-gatherer subsistence, one site-specific and the other regional in scale. At the site-specific level, we proposed to focus on the materials from heavy fraction samples derived from the flotation of feature matrix.

In addition to this site-specific data, to place the site within a regional context, we proposed to collect data on subsistence

resources present within the assemblages of the sites that were also used in the investigation of regional land use patterns. If and when available, we proposed to collect data on the following aspects of the assemblages: (a) proportion of marine and terrestrial fauna represented in the assemblage; (b) taxa present by prey size classes (small, medium and large mammals); (c) distribution of taxa by habitat within each assemblage; (d) if feasible, we proposed to investigate the potential for field processing of resources by determining the skeletal elements represented at each site as well as the mollusc species present and their caloric yield.

Available Data and Analyses Conducted: Seventy-one heavy fraction samples were used for this analysis. They consisted of the 66 samples submitted for macrobotanical analysis and five other float samples not from feature contexts. The mollusc samples derived from the heavy fractions were sorted into species. Animal bone was analyzed to determine the species represented. Unfortunately, the bulk of the sample was unidentifiable to the taxa level. The small amount of mammalian bone recovered from 41NU2 was not quantified by size category into very large, large, medium, small and very small groups as was proposed in the research design. Also, because none of the mammalian bone could be identified to element, the meat utility of the skeletal elements was not determined. Finally, while in the original research design we proposed to submit for macrobotanical analysis all charcoal samples that were not needed for paired sample analysis, this was not done. Instead, we limited the macrobotanical analysis to heavy fraction samples obtained from well provenienced feature-associated contexts (i.e., the 66 samples mentioned above).

To address the regional-level research objectives, we collected the data proposed in the research design from the same 18 sites used in the regional land use investigation. The 18 sites and/or components were: 41NU11, 41NU184, 41NU221, 41NU266, 41NU267, 41NU268, 41NU276, 41SP15, 41SP43, 41SP120, 41SP136, 41SP148, 41SP149, 41SP153, 41SP156, 41SP177, 41AS3 and 41AS16. Some data types that were originally proposed (i.e., breakdown of skeletal elements) could not be collected due to lack of information or necessary detail in the reports.

Research Problem: *Relationship Between Distribution of Cemetery Sites and Territoriality*

Given the prevalence of cemetery sites and sites with human burials within the Coastal Plains and Marshes region, and given the cemetery that has been documented at 41NU2, a better understanding of the role of cemetery sites may improve our understanding of the function of 41NU2.

Associated Research Questions:

- (1) What is the distribution of sites with burials across the Gulf Coastal Plains and Marshes region?
- (2) What are the ages of the cemetery components or burials present at each site?
- (3) What are the habitats that contain sites with burials?
- (4) Are sites with individual or small numbers of burials distributed differently in space and time than sites with recognized cemeteries?
- (5) Are there differences in burial practices between sites with small numbers of interments versus sites with cemeteries?
- (6) Are there any indicators of status differentiation between individual burials in cemetery sites?
- (7) What is the incidence of violent deaths in cemetery sites and is there any temporal or spatial patterning in the data?

Proposed Research: In the project research design, we proposed to create a comprehensive Geographic Information System database containing information on all sites with human remains identified throughout the Gulf Coastal Plains and Marshes and immediately surrounding regions (e.g., South Texas; South-central Texas). This database was to include information on individual interments at cemetery sites (e.g., age, sex, orientation, associated grave goods, association with other individual). Additionally, we proposed to collect information on the ages of cemetery-associated components, site size, habitat where site is found; and whether the site had a habitation component. The database was to be employed in spatial analysis to determine whether there are any trends in the spatial distribution of burial-associated sites and whether there are any temporal trends in the numbers and/ or types in associated grave goods and the incidences of violent deaths.

Available Data and Analyses Conducted: The CAR monitoring project identified and exhumed six human remains and Ricklis (1997) also found one. In addition to this small sample, CAR also obtained information on an additional 103 individuals buried at the site and later exhumed by Jackson (see Jackson et al. 2004). These data were added to the regional database. The regional-scale comparative database was compiled from the existing literature on burial sites distributed across the Gulf Coastal Plains and Marshes. We relied on a number of very comprehensive recent summaries compiled by Hall (1995b, 1998) and Perttula (2001) and augmented them by a thorough search of the Texas Archeological Sites Atlas.

Research Problem: Completion of the Inventory of Human Remains for NAGPRA Purposes

One of the main tasks associated with this project was the completion of the inventory of human remains recovered by CAR during its road construction monitoring activities. This was important because of TxDOT-ENV's obligations regarding the Native American Graves Protection and Repatriation Act.

Proposed Research: In the research design CAR proposed to conduct a 100 percent inventory of human remains from seven individuals uncovered and returned to the CAR laboratory during the project.

Available Data and Analyses Conducted: The complete this requirement, standard osteometric analyses were conducted on the human remains recovered by CAR. The analyses were conducted in accordance with the criteria set forth in the Standards (Buikstra and Ubelaker 1994). The forensic data was compiled for and evaluated in the Forensic Database (ForDisc v.2.0) maintained by the University of Tennessee at Knoxville (Moore-Jansen et al. 1994). The staff human osteologist/physical anthropologist carried out these analyses. The inventory was submitted to TxDOT-ENV under a separate cover.

Chapter 10: Paleoclimate

Steve A. Tomka and Raymond P. Mauldin

Previous Reconstructions Of Paleoclimate

The data on the modern climate described in Chapter 2 suggest that the climate currently associated with 41NU2 reflects a relatively warm environment with variable monthly and yearly precipitation. However, the degree to which these general climate characteristics are reflective of the past is unclear. While research, particularly during the past decade, has contributed immensely toward understanding the paleoenvironment of the state (e.g., Bousman 1998; Caran 1998; Cooke 2005; Frederick 1998; Fredlund et al. 1998; Nordt et al. 1994; 2002), much of that research has focused on Central, South-Central, and Northern Texas. With a few exceptions, most of which are primarily concerned with shifts in sea levels during the Holocene (e.g., Albert 2006; Blum et al. 2002; Ricklis 2004; Ricklis and Blum 1997; Ricklis and Cox 1998), southern and especially gulf coastal areas have not been extensively studied, a phenomenon clearly tied to a lack of appropriate data sets in these areas of the state. Consequently, what we currently think we know about paleoenvironments in coastal and near coastal settings of Texas is primarily dependent on studies in other portions of the state, as well as on phytolith research conducted along the coast roughly 25 years ago (Robinson 1979, 1982).

Recent overviews of paleoclimate in other portions of the state are available (e.g., Bousman 1998; Cooke 2005; Nickels and Mauldin 2001; Nordt et al. 2002). Figure 10-1 presents a comparison of several of these climate summaries derived from 1) soil organic carbonate from south-central Texas (ca. 200 km to the north of 41NU2), 2) carbon isotope measurements of humic deposits in Hall's cave (ca. 340 km to the northwest of 41NU2), and 3) arboreal and grass pollen from bog deposits (ca. 300 km north of 41NU2). These data sets should all be responding to vegetation changes, though the precise temporal and spatial scale of the shifts in climate variables relative to shifts in these climate proxy data sets is unclear. While these sequences differ in some details, the general pattern that emerges from these (Figure 10-1), as well as earlier studies, is one in which the Early Holocene (ca. 10,000-8,000 B.P.) reflects a cooler, mesic climate, with a slight, gradual drying trend reflected in some data sets throughout the 2,000 year period. A warm and dry Middle Holocene (ca. 8,000-4,000 B.P.) follows, perhaps with a brief mesic period sometime around 7,000 to 6,000 or 5,000 B.P. A return to xeric conditions between 5,000 and 4,000 B.P. follows that mesic interval. The Late Holocene (ca. 4,000 to 0 B.P.) is thought to be generally cooler and/or wetter, with

a return to more xeric conditions within the last few hundred years, though this is beyond the scope of the data set.

Coastal climate research is difficult to interpret in light of these more general, statewide patterns. For example, Robinson (1979) discusses the opal phytolith records from two sites (41GD21, 41GD21A) on Colette Creek in South Texas (Robinson 1979:111) that are located roughly 100 km to the north of 41NU2. Eight samples were processed from six different contexts on these sites. Like those data sets shown in Figure 10-1, shifts in opal phytoliths are reflective of shifts in vegetation. Unfortunately, dating of most of these sample contexts is problematic. Samples 6 and 5 can only be assigned dates of older than about 4,450 B.P., while Sample 4 is associated with a $c14$ date of 4,260 \pm 250, a date that translates into a roughly 900 year range (5,300- 4,400 B.P.) at 1-sigma. Sample 3, again based on radiocarbon dates, can be placed between 5,300-4,400 B.P. and 3,350-2,300 B.P., while two samples (designated sample 2, 2a), can be placed sometime within a 1,050 year period (3,350-2,300 B.P.). Finally two samples designated sample 1 and 1a are associated with Late Prehistoric material, though that material is also mixed with historic remains. Robinson's data (1979), then, suggest that sometime before about 4,850 B.P., phytoliths from the deepest sample (Sample 6) are reflective of a well developed forest in this region. A more recent sample, but one that also dates sometime before 4,850 B.P., produced phytoliths that suggest an extremely dry period was present. Sometime between 5,300 and 4,400 B.P., there appears to have been a brief mesic period, followed by a dry climate sometime between 4,850 and 3,350 B.P. At around 2,825 B.P. (samples 2, 2a), phytoliths suggest a mesic interval, with a more xeric mixture of both wet and dry species present during the late prehistoric or historic periods.

Additional phytolith work has been conducted by Robinson (1982) at Choke Canyon on sites located roughly 110 km to the northwest of 41NU2. Using a variety of samples, Robinson (1982:598) infers a cool, mesic climatic regime that shifts to a more arid period sometime prior to 5,300 B.P., a pattern that continued between about 5,300 and 4,300 B.P. He then suggests a return to both cooler and wetter conditions by 3,250 B.P. The Choke Canyon analysis points to mesic conditions (similar to today's) by about 2,540 B.P. (Robinson 1982:598-599). Afterward, a shift to more xeric conditions occurred by about 1,000 B.P., but Robinson suggests that they may have been more mesic than modern conditions (Robinson 1982:599). Unfortunately, like the Colleto Creek

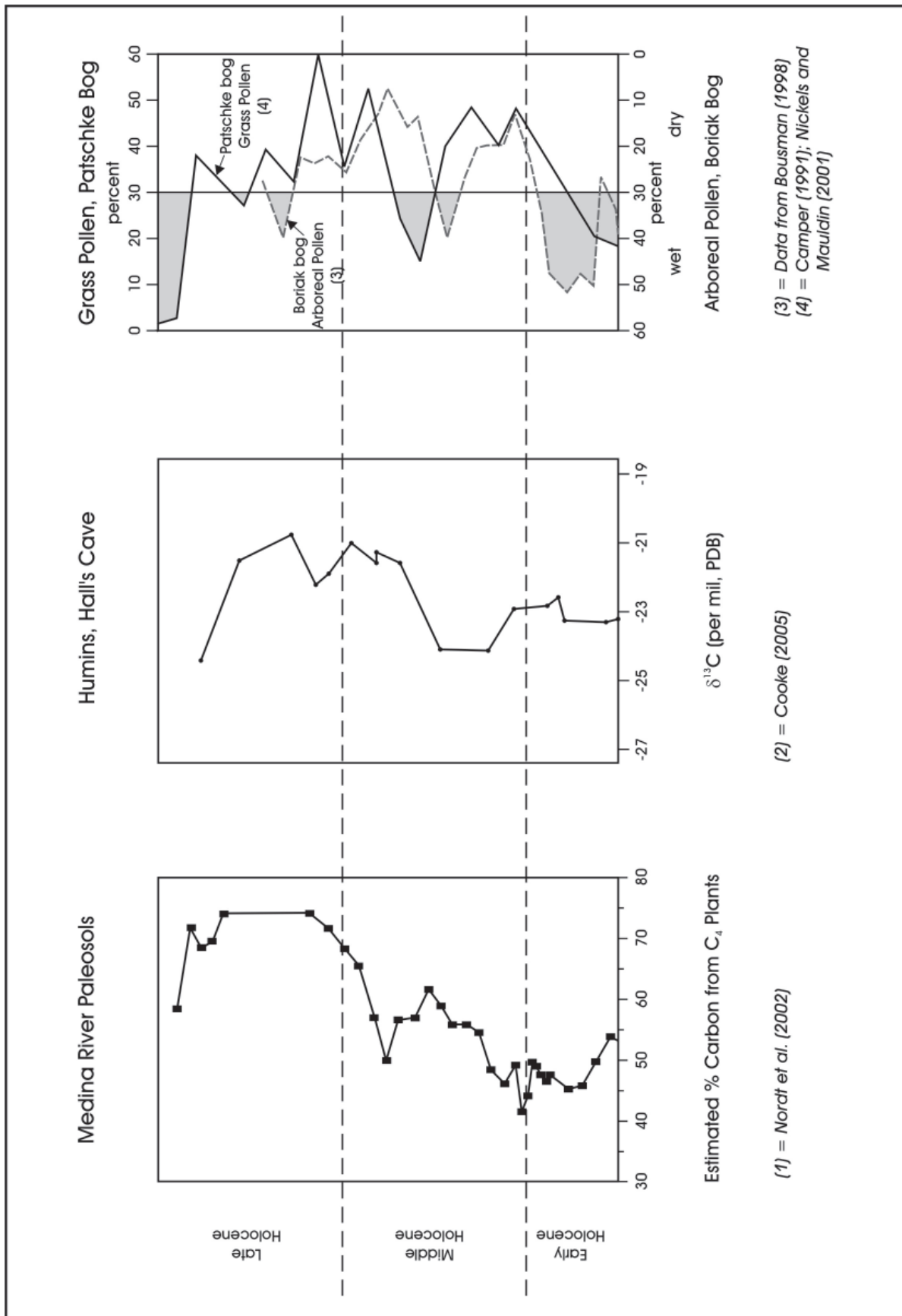


Figure 10-1. Comparison of climatic sequences.

work discussed previously, the samples from Choke Canyon are poorly dated and underreported. It is difficult, therefore, to rely heavily on Robinson's work.

K.A. Cox and S.A. Cox (1993) and Ricklis and Cox (1998) present a study of oyster habitat characteristics and growth patterns from a series of sites in the Middle Coast. They suggest that estuarine salinity, measured by reference to an ordinal scale developed from the presence and characteristics of oysters, was highest between about 5,000 and 3,700 years ago, with declining salinity after 3,700 B.P. They suggest that these patterns may be a result of low freshwater discharge between 5 and 3.7 Kya, and increasing freshwater runoff, probably reflecting increasing rainfall, after 3,700 B.P. K.A. Cox and S.A. Cox (1993:94-95) and Ricklis and Cox (1998:129) suggest that these patterns may be indicative of a dry Middle Holocene, with increasing moisture present in the Late Holocene. Ricklis and Cox (1998) also use the percentage of oyster shells from an assemblage that have "readable" winter seasonality indicators in growth rings. Oyster assemblages with high "readability" were probably formed during cooler periods, while growth of oysters in assemblages dominated by low readability probably occurred during periods of warmer temperatures. Assemblages with high readability percentages, and by extension cooler temperatures, occur at around 7,200 B.P., 5,100 B.P., and between about 2,750 and 1,750 B.P. Conversely, warmer temperatures are indicated from about 4,600 to about 3,000 B.P. (Ricklis and Cox 1998:132). While the connections are tentative, if changes in oyster habitat and seasonality indicate differing rainfall and temperature regimes, the patterns presented by Ricklis and Cox (1998) are consistent with the general patterns outlined previously in which a mesic Early Holocene yields to a xeric Middle Holocene, possibly with a brief mesic interval sometime around 5 to 6 Kya. An initial mesic interval in the Late Archaic follows the warm, dry Middle Holocene.

Coastal Occupation Patterns and Paleoenvironmental Conditions

Additional paleoenvironmental research, primarily concerned with documenting Holocene sea level changes and proposed development of coastal landscapes, has been conducted in many coastal settings, including the Texas Gulf Coast (e.g., Albert 2006; Anderson and Thomas 1991; Curray 1960; Nelson and Bay 1970; Wilkinson and Basse 1978). In the current context, these Holocene sea level shifts are of considerable interest as Ricklis (1995, 2004) has suggested that prehistoric occupation patterns of coastal areas in Texas are closely tied to shoreline shifts. In his view, periods of sea level standstill provided opportunities for the evolution of lagoon/estuary systems along the Texas coast that are highly

productive settings for hunters and gatherers. Conversely, periods of rapid sea level rise would have damaged these aquatic ecosystems through increased salinity and disruption of nutrient rich shallows (see discussion in Ricklis 1995, 2004; Ricklis and Blum 1997).

Figure 10-2 is based on calibrated, corrected radiocarbon dates taken from shell and charcoal samples presented in Ricklis and Blum (1997:Table 1). These dates are primarily from archeological sites in Nueces, San Patrice, Refugio, and Calhoun Counties of the Middle Coast. The plot was constructed by using 78 of the 80 dates; two dates with large sigmas were eliminated. The 78 dates were then calibrated using OxCal (Ramsey 2003), and plotted at a 2-sigma range. These plots were then converted into a line graph with an interval of 200 years. One count was made for each time that a date range fell within a given 200 year period. To the degree that the resulting plots reflect Middle Coast occupation patterns, the data clearly suggest two periods of absence, with the most recent occurring between about 3,800 and 3,400 B.P., and the other occurring between 6,700 and 6,000 B.P. Ricklis (1995; 2004; Ricklis and Weinstein 2005; see also Ricklis and Blum 1997) correlate these two absence periods with periods of rapid sea level rises, and periods of occupation with periods of sea level standstill or slow sea level rise. He concludes that by 3,000 B.P., sea levels stabilized near their modern position, allowing development of the modern barrier islands that protect the lagoons and bays and decreased the salinity of the bay waters, resulting in a dramatic increase in aquatic productivity. Using a variety of measures of occupational intensity, Ricklis (1995; 2004) further shows that a dramatic increase in the intensity of human use of shorelines does occur after 2,000 B.P., an increase consistent with the model. The Figure 10-2 data also suggest increased occupational intensity after around 2,000 B.P.

While it is clearly the case that occupational intensity does vary in this section of the coast, and while the patterns presented by Ricklis are striking, the proposed explanation linking varying rates of sea level increases and periods of standstill with high or low estuary productivity, and associated occupation to take advantage of that productivity, appear to rely exclusively on one of several competing models of sea level changes. In fact, most recent studies of sea level curves for the Gulf coast, as well as data from Bermuda, Barbados, and Belize see Gischler and Hudson 2004; Mackinnon and Jones 2001; Parkinson and White 1994) suggest a model quite different from that used by Ricklis. These more recent models note relatively rapid, continuous and early rise in sea levels up to sometime around 5,000 to 6,000 years ago, with slow increases after that point (see Blum et al. 2002; Otvos 2004; Törnqvist et al. 2004). That is, these models appear to eliminate any prolonged periods of standstills, especially prior to about 2,000 to 3,000 years ago

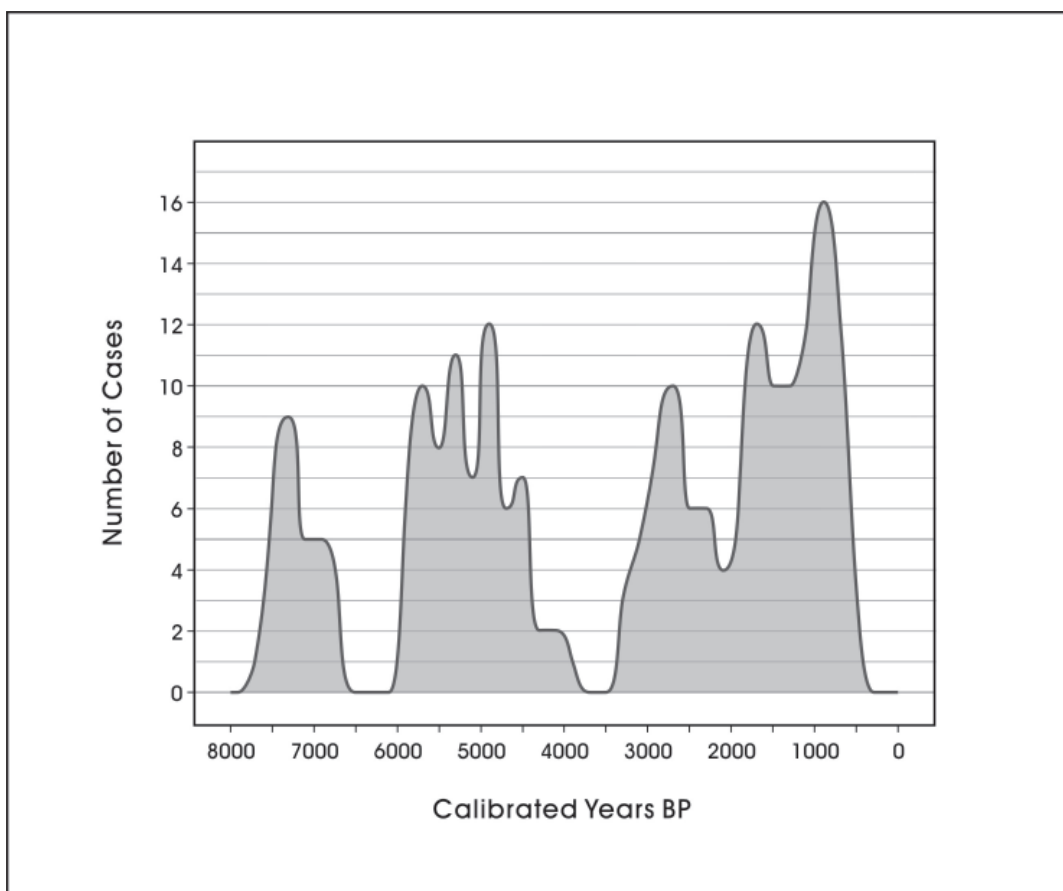


Figure 10-2. Corrected radiocarbon dates from shell and charcoal samples reflecting middle coast occupation.

when essentially modern shore lines may be achieved. They also eliminate any dramatic, rapid alterations in sea level rise that would produce the ecological consequences suggest by Ricklis to have caused abandonment of the Middle Coast at various points.

A recent study of pollen from soil cores taken at the Swan Lake site (41AS5) in Aransas County (Albert 2006) indicates a vegetation history that, at first glance, seems to support Ricklis' (1995; 2004; Ricklis and Weinstein 2005; see also Ricklis and Blum 1997) model. Albert (2006) found that the pollen in the soil cores, radiocarbon dated to approximately the last 7000 years, indicated fluctuations in plant communities forming 6 "pollen zones" (Albert 2006:166). Pollen Zones 1 (7100-6650 B.P.), 3 (4900-3850 B.P.), and 5 (ca. 3450 B.P.) have a high percentage of pollen from the Chenopodiaceae family. Pollen Zones 2 (5800-5400 B.P.), 4 (ca. 3650 B.P.) and 6 (3050 to the present) have low percentages of Chenopodiaceae and higher percentages of oak (*Quercus*), and various grass and forb pollen (Albert 2006:166-172). He interprets these patterns as indicative of changing sea levels in which the wetlands at 41AS5 were invaded by rising sea

levels during Zones 1, 3, and 5, turning the area in to a salt marsh, and reducing overall productivity. He believes the higher percentage of arboreal and forb pollen in pollen Zones 2, 4, and 6 indicate a period of sea level stand still, during which the area is a freshwater wetland and the productivity increases substantially (Albert 2006). Albert (2006:173-174) suggests that periods in which the pollen data indicate that the Swan Lake area was a salt marsh roughly correlate with Ricklis' estimations of periods during which the coast populations were low or non-existent. However, Ricklis (1995:277) identified two periods of abandonment, one between 6800 and 5800 B.P., and a second between 4200 and 3100 B.P. (see also Figure 3-10). The initial abandonment does have a significant overlaps with Albert's Zones 1 (7100-6650). The second period of abandonment (4200-3100 B.P.), however, overlaps with Zones 3, 4, and 5- (ca. 4900-3450), and encompasses periods of both low and high productivity.

While, on the whole, modern climate conditions along the coast are wetter and warmer than the rest of the state, the area associated with 41NU2, located in the southern section of the Middle Coast, is warm and dry, with high variability in

precipitation. The degree to which these modern patterns were altered by changing climate regimes over the last 10,000 years is not clear as there are significant gaps in our understanding of critical paleoclimate parameters in the coastal region. While innovative phytolith research along the coast, conducted by Robinson (1979, 1982), suggests that significant alterations in vegetation structure has occurred over time, the studies lack good temporal control and are difficult to interpret. Preliminary work on oyster growth patterns and measures of salinity conducted by K.A. Cox and S.A. Cox (1993) and Ricklis and Cox (1995) near 41NU2 seem to be consistent with other, better controlled paleoclimate investigations that suggest a warm and dry Middle Holocene and a more mesic Late Holocene. These poorly documented shifts in climate, and associated vegetation shifts, should have produced differing mixes of plants and animals for exploitation by hunters and gatherers. An overview of occupation on the coast using radiocarbon dates, as well as volumetric comparisons of archeological materials from coastal sites conducted by Ricklis (1995, 2004; Ricklis and Blum 1997), suggest that the coast was used, at varying levels of intensity, over much of the Middle and Late Holocene, with increasing use present after about 2,000 B.P. A model (see Ricklis and Blum 1997) linking occupational intensity with varying rates of sea level increases and periods of sea level standstill, however, uses Holocene sea level curves that have recently been questioned by a number of researchers. While additional work on this problem is clearly necessary, the most recent models of sea level changes in the Texas Gulf do not seem to support the stepped pattern suggested by earlier researchers. That is, shifts in productivity do not follow the most recent sea level curves. Of course, a variety of factors other than major shifts in sea level could impact estuary productivity. These factors include shifts in the frequency and intensity of storms, the formation or destruction of barrier islands, and inland shifts in rainfall and runoff. Clearly, additional work, along the lines undertaken by Albert (2006) is warranted to explore Ricklis' model.

The Data from 41NU2

During August 2007, six cores were extracted from the ROW of Ennis Joslin and within the boundaries of 41NU2 (Figures 5-1, 5-2, 5-4). These cores were inspected in the field to determine the amount of intact deposits they sampled. Of the six cores, three contained intact stratigraphic column and were subsequently sub-sampled at the CAR laboratory. The sub-samples included matrix for pollen and phytolith analyses, ostracod and mollusc analysis, diatom analysis and lithological analyses. Core 2, was extracted from within the limits of a pond deposit previously identified near the intersection of Ennis Joslin and Ocean Drive. Core 4 was recovered from the east ROW of the north dune near station 34+050. Core 6 was recovered from the east ROW of Ennis Joslin near station 34+550.

The remaining portion of the chapter pulls together the interpretations derived from the specialized analyses and compliments this with the paleoenvironmental implications of the data collected from the analysis of the molluscs and fish otholiths. We begin with the consideration of the results of each core before we shift to the discussion of the entire site.

Core 2

Cummings et al., (Appendix A, this volume) divide the pollen record of this core into five zones (E-A). Poaceae, Chenopodium and High-spine Ateraceae are the three co-dominants representing grasses, members of the goosefoot families and likely amaranth, and a large group of plants that include the sunflower family. Significant shifts in microscopic charcoal relative to quantities of pollen are noted throughout the record. Both the pollen and phytolith records show some subtle and other long-term environmental changes as well as punctuated events and/or short term changes detectable even at the course-scale sampling strategy employed. The Zone E pollen record contains alder and birch suggesting their regional presence. Bryant (Appendix B, this volume) suggests that the identification of these species may be mistaken due in part to the degraded condition of the pollen. Another possible explanation comes from Frederick's (Appendix F, this volume) findings. His stable carbon isotope analysis of Core 6 from the southern end of the site suggests that at the bottom of the stratigraphic profile, dated to around 4000 yrs B.P., a woodland environment may have been present. He suggests, however, that this stable carbon isotopic signature may be the product of the deflation of the former Pleistocene surface. Getting back to the Core 2 record, the presence of small quantities of Pinus and Quercus pollen within Zone E of Core 2 may suggest that these species have been present in the regional vegetation since at least 3,000 years ago. The vegetation community may have consisted of mesic habitats, woodland margins and freshwater marsh habitats. The phytolith record from this zone appears to be indicative of a coastal freshwater marsh/pond lake/stream habitat. Forest edge and understory grasses are very low in occurrence while sandy soil grass and shoreline and emergent aquatic grasses are fairly high in abundance. Overall, the pollen and phytolith record from Zone E represents an expansive freshwater marsh habitat with dunes and tidal mudflats extending beyond their modern limits likely due to the somewhat lower sea levels (approx., 1.0 meters; see Flemming et al. 1998; Milne et al. 2005).

The pollen record from Zone D contains indicators of beach and grasslands and thickets. The presence of *Persicaria* pollen suggests periodic changes in water levels or wave action that creates a linear continually disturbed zone along the shore. The phytolith record from the zone supports the suggestion of favorable coastal freshwater conditions. The recovery

of freshwater sponges may correspond with increased precipitation, a rising water table, or potentially less saline conditions in ponds and the surrounding environment. The slight increase in the bilobate phytolith from forest edge and understory grasses also supports the interpretation that precipitation may be on the increase. The only zone in which *Isoetes* (quillwort) pollen is present. The presence of the pollen suggests cooler than modern temperatures and the combined pollen and phytolith record suggests cooler and wetter conditions than current conditions (perhaps even stormier).

The Zone C pollen assemblage (predates 1300+/-40 RCYBP comes from the base of Zone B) suggest the expansion of the freshwater marshes. The frequencies of bilobate indicators of grasses that prefer sandy substrates is at an abundance in contrast to the forest edge and scrubland grass phytolith. The patterns may correspond with increases in storm frequencies during the depositions of Zones C and D. This is supported by the fact that Zones C and D exhibit the highest rates of sedimentation (10.6 years/cm).

The base of the Zone B dates to 1300+/-40 RCYBP. The pollen assemblage is consistent with a drying trend. The phytolith record indicates the development of slightly brackish marsh conditions. Bilobate grass taxa associated with forest edge scrubland ecotones reach their peak in this zone. Phytolith assemblage is also suggestive of more brackish marsh conditions potentially coupled with warmer and drier conditions that I the underlying zone.

The top of the stratigraphic column in Zone A is dated to 550+/-40 RCYBP. Its pollen assemblage is characterized by an increase in *Quercus*, *Cheno-am* and *Cyperaceae* pollen and the appearance of *Prosopis* pollen. The phytolith assemblage suggests a decrease in the extent of the brackish marsh habitat and an increase in the drier habitats such as sand dunes in the environment. Phytolith from dry habitat grasses that prefer sandy soil substrate increase in frequency. The zone also contains the highest levels of charcoal and oak in the pollen record.

Winsborough's (Appendix C, this volume) analysis of the diatom record from Core 2 indicates that the pond was likely a shallow body of water sustaining a somewhat polluted freshwater habitat with small to moderate levels of salt. The substrate appears to have been somewhat muddy and water levels fluctuated seasonally suggesting that freshwater was only available on a seasonal basis and during the warm dry season (summer) the pond may have drier out entirely.

Palacios-Fest's (Appendix E, this volume) analyses of Core 2 samples contained no marine ostracods. The foraminifera and

mollusc records were poor from Core 2 though the patterns are consistent with the ostracod record indicating that the pond was a source of fresh water throughout its existence. The fauna of Core 2 may have been transported into the pond. Particle size analysis conducted by Palacios-Fest as part of the extraction of the faunal record suggests that overall the inflow of water into the pond was slow. The sediment particle analysis indicated that there were two episodes of rapid influx based on the coarsening of the sediments, one between 2970 and approximately 2500 B.P., and the latest between 700 yrs. B.P. and the present. Two episodes of slower discharge are also suggested by the particle size analysis. The first occurs around 2500 B.P. following the rapid discharge and the second around 1000 B.P., prior to the rapid discharge episode mentioned above. It is difficult to determine whether these coarse and fine deposits reflect different depositional regimes (i.e., slope wash from the dunes or eolian sedimentation from the nearby mudflat or differences in the sources of the parent materials being transported by wind onto the dunes (Frederick 2008, Personal Communication). The Core 2 paleosalinity pattern is only available for the later part of the chronological sequence and for the period between 1300 and 550 yrs B.P. it shows relatively low salinity brackish conditions. Again, this pattern may be derived from the decreased salinity of the source deposits in the mudflat or an increase in the influx of fresh water due to increased precipitation in the region.

Frederick's lithological analysis (Appendix F, this volume) of Core 2 samples indicates a pattern similar to that noted by Palacios-Fest. The percentage of sand is high and steadily decreasing between 2970+/-40 B.P. and 2240+/-40 B.P. The proportion of sand-sized particles remains relatively low with occasional small increases only until around 1000 yrs B.P. at which point it again increases until the present.

Core 4

Cummings et al. (Appendix A, this volume) note that the Core 4 pollen and phytolith record has been severely affected by silica dissolution. One sample (#182) recovered from 40-43 cmbs and dated to approximately 500 RCYBP, is equivalent to Zone A or above in Core 2. The phytolith record from this sample suggests a sand dune or wet meadow interdune environment with a relatively brackish and shallow watertable. Grasses that were well adapted to sandy substrate dominated the vegetation.

The diatom assemblage from Core 4 is very small and is likely the result of wind transport from the nearby lagoon (Winsborough Appendix C, this volume). The small assemblage confirms that the water level and the edge of the lagoon shifted over time at times getting closer to the sample locality while at other times receding to some distance. The

assemblage also suggests that the lagoon may have oscillated either seasonally or for extended time, between containing water, becoming muddy and even entirely drying out.

The Core 4 samples contained a variety of ostracods mainly of marine origin most likely deposited through wave action, according to Palacios-Fest. While the edge of the bay could have been sufficiently close to influence the edge of the dune, it is more likely, that eolian deposition is responsible for the presence of these marine ostracods in the dune sediments. The foraminifera record from the core was consistent with the ostracod assemblage. On the other hand, the mollusc record was poor from Core 4 but showed that only marine species occurred in samples from this core. Again, these molluscs were likely present in the source materials that were eroded through wind action to contribute to the building of the dune. Palacios-Fest's particle size analysis from Core 4 shows three periods of sediment coarsening. The oldest occurs around 3620 B.P., the second around 2110 B.P. and the most recent sometime after 1060 B.P. The period between the first and second coarse sediment deposits appears to have been characterized by slow fine sediment deposition resulting in the accumulation of primarily clayey sediments. Palacios-Fest suggests that the fauna recovered from Core 4 indicate that the periods of coarse sediment deposition were the result of freshwater inflows. However, given Frederick's conclusion that the sediments composing the dune are eolian in origin, it is more likely that the coarsening of the deposits is due to a change in the character of the parent materials that contributed sediments to the dune. The clayey sediments are also the product of eolian deposition and again may reflect a fining of the parent materials contributing sediment to the dune. The paleosalinity data supports this interpretation.

Frederick's lithological analysis (Appendix F, this volume) of Core 4 samples indicates that sand-sized particles were most common prior to 3620 \pm 40 yrs B.P. and that the sediments began to become finer likely beginning around 4000 yrs B.P. Between 3620 \pm 40 yrs B.P. and 2110 \pm 40 yrs B.P. sediments tend to remain clay dominated although there is a small peak in sand-sized particles around 2880 \pm 40 yrs B.P. Sand-sized particles again increase in proportion between 2110 \pm 40 B.P. only to begin declining before 1060 \pm 40 B.P.

Core 6

Only four of the 16 samples examined by Cummings et al. (Appendix A, this volume) from Core 6 yielded statistically significant phytolith counts. The deepest of the phytolith assemblages comes from Zone G which pre-dates 3770 \pm 40 RCYBP. The phytolith assemblage is unlike any modern Gulf Coast habitat assemblage. Phytolith diagnostic to the

palm family were the most abundant form and may derive from dwarf palmetto or Rio Grande palmetto. The presence of this phytolith suggests that a maritime forest of freshwater swamp with an arboreal canopy was the dominant vegetation community around 4000 RCYBP. The phytolith assemblage from Zone F of the core is severely affected by silica dissolution. Keeping in mind that the record that remains is heavily biased toward certain more resistant morphotypes, the assemblage continues to indicate the presence of a forested canopy with a palm component around 3770 \pm 40 RCYBP. The base of the core has reached the top of the Beaumont Clay and therefore it is possibility that deflated Pleistocene phytoliths are being sampled by this zone.

Cummings et al. (Appendix A, this volume) were able to define a Zone E phytolith assemblage from the Core 6 samples. The severely affected assemblage dates to around 2760 \pm 40 RCYBP and appears to be roughly contemporaneous with Zone E in Core 2. The degraded phytolith assemblage seems to suggest that by this time, the vegetation community has changed to an expansive freshwater marsh. The surviving phytolith sample from a soil sample that is equivalent to Zone D from Core 2, suggests the presence of a sand dune habitat with a shallow watertable at around 1770 \pm 40 RCYBP. Samples from what would be equivalent to Zone C in Core 2 are also heavily affected by silica dissolution. The biased record suggests the existence of moist sand dune/interdune conditions. Samples for what would be equivalent to Zone B in Core 2 are radiocarbon dated to 1280 \pm 40 RCYBP. The affected partial record suggests that a salt marsh existed in the area around that time. The best preserved phytolith assemblage comes from Zone A which is roughly contemporaneous with Zone A from Core 2. The record suggests the presence of a dune habitat with elevated watertable. Moist interdune meadows may have been present nearby. In combination, the Zone B and A records suggest a pattern of increased drying beginning during the upper levels of Zone B and proceeding through Zone A and perhaps representing the shift from a salt marsh to a drier sand dune environment. This trend is analogous to that noted in the equivalent zones of Core 2 and potentially contrary to the stable carbon isotope data.

The diatom assemblage from Core 6, analyzed by Winsborough (Appendix D, this volume), contrasts with that of Core 2. The Core 6 assemblage reflects the sandy substrate that characterized the edge of the dune from where the core samples originated. The species reflect seasonally wet conditions and the absence of muddy substrate species indicates that water accumulations were short lived and may have occurred in clay-lined depressions or swales in the sand.

Core 6, similar to Core 4, contained a variety of ostracods mainly of marine origin most likely deposited through eolian

transport although wave action cannot be entirely ruled out. Core 6 contained a varied foraminifera record consistent with the ostracod assemblages. The mollusc record was poor but only marine species occurred in the samples. The fauna from Core 6 indicates that the sediments sampled by the core may have received a relatively constant input of freshwater. This is consistent with the location of the core on the bank of a small arroyo that probably carried water during rain events and wet periods. The faunal record does indicate that the largest pulses of freshwater came between 3770 and 2760 B.P. The strongest seawater influence appears to have occurred sometime just after 1770 \pm 40 yrs B.P. perhaps representing a dry period characterized by an increase in eolian deposition. In general, the Core 6 paleosalinity pattern suggests that this location was much more susceptible to freshwater influxes than the locality adjacent to Core 4.

Palacios-Fest's particle size analysis of the sediments from this core show an increase in sand-sized particles around 2760 \pm 40 B.P. A period of clayey particle deposition occurs between about 1770 \pm 40 B.P. and 1280 \pm 40 B.P. which is followed by a peak in coarse particle deposition. The most evident increase in coarse particles in the Core 6 data in Frederick's analysis occurs between about 2760 \pm 40 B.P. and about 1770 \pm 40 B.P. Between 1760 \pm 40 B.P. and 1280 \pm 40 B.P. sediments tend to consist of clayey particles. This period is followed by a period of increase in particle size. The results of the two particle size analyses are very similar.

The Combined Picture: Patterns and Disagreements

Based on the analysis of the pollen and phytolith samples from the three cores from 41NU2, Cummings et al. (Appendix A, this volume) define seven climatic, habitat and landscape-level changes at the site. The oldest, dating to around 4000 yrs B.P. is manifested in Core 6 and suggests the presence of a maritime forest or fresh water swamp with an arboreal canopy flourishing under mesic and warm conditions. This forested swamp community appears to have persisted until at least 3770 \pm 40 B.P. under wet but possibly cooler conditions that may have resulted in the thinning of the arboreal canopy. By 2970 \pm 40 yrs B.P., the freshwater pond just north of 41NU2 begins to accumulate water and may be at the center of a sedge or freshwater marsh. The forested swamp at the south end of the site continues to thin and gradually changes to a relatively open freshwater marsh by around 2670 \pm 40 yrs B.P. Between 2240 \pm 40 yrs B.P. and 1280 \pm 40 yrs B.P. the water level within the pond at the north end of the site fluctuates and becomes increasingly shallower and more brackish and the marsh at the south end becomes a salt marsh. Rainfall appears to increase during this period while

temperatures may become cooler. After about 1280 \pm 40 B.P. a trend that may have begun as early as 1800 \pm 40 B.P. sees the increase in oak and cattail pollen. Sand dunes are well established by 550 \pm 40 B.P. and the pond supports a brackish marsh and moist interdune areas may be present as temperatures appear to increase and precipitation decreases leading to the formation of the scrublands community characterizing the area today.

Overall, the diatom assemblages from the three cores (Winsborough Appendix D, this volume) indicate that rain was episodic and seasonal resulting in temporary accumulations of fresh water in the pond. However, it is likely that these accumulations did not last through the dry season. These conditions suggests that unless fresh water was transported into the locality, its seasonal availability would have favored and limited the occupation of the site to the wet season or season, namely Spring and Fall and following episodic occurrences brought on by hurricanes.

According to Palacios-Fest (Appendix E, this volume), the paleosalinity patterns in combination with the ostracod, foraminifera and mollusk data allow the definition of three paleoecological zones characterized by more or less saline conditions. The reconstruction offered by these data types is quite different from that based on the pollen and phytoliths offered by Cummings et al. (Appendix A, this volume). High salinity, reflecting xeric conditions appears to have dominated the area before 3600 years B.P. The interval between 3600 and 1300 B.P. appears to have been one of transition from xeric to more mesic conditions while the last 1300 years are characterized by relatively mesic conditions. Unfortunately, much of this reconstruction is derived from Cores 4 and 6 and even within them, substantial variation is present that is related to their location in relationship to topographic features such as crest of dune, distance to nearby arroyo, etc.

Frederick's analysis of delta 13 carbon from Cores 4 and 6 from the north and south dunes respectively, provides an additional composite image of paleoenvironmental conditions in the vicinity of 41NU2. His analysis indicates that the vegetation community characterizing the site vicinity was more mesic adapted around 4000 yrs B.P. However, between about 3770 \pm 40 yrs B.P. and 1000 B.P. there was a steady movement in the composition of the vegetation community to a greater and greater proportion of arid-adapted species. A slight shift back to more mesic-adapted vegetation seems to characterize the last 1000 years.

Overall these diachronic trends appear to correspond well with the reconstructions based on pollen and phytoliths offered by Cummings et al. (Appendix A, this volume).

The maritime forest or arboreal community surrounding a freshwater swamp present near the site around 4000 B.P. gradually becomes more open over time giving way to sand dunes with moist interdune settings as precipitation decreases. The return of the woody vegetation during the last 1000 years of the sequence may reflect a combination of increased moisture and historic impacts on vegetation communities. The findings by Dering (Appendix C, this volume) of mesquite (*Prosopis glandulosa*), carbonized fragments of a woody legume (*Prosopis* or *Acacia* sp.), and oak (*Quercus* sp.), coupled with the oak pollen identified by Cummings et al. (Appendix A, this volume) support the trends in the stable carbon isotope data. Dering identified mesquite in five features (Features 102, 114, 156, 165, and 168) and woody legume fragments in fourteen features. None of the features containing mesquite have been dated but two of the woody legume containing-features have dates. Feature 121 returned a date of 2740 \pm 40 B.P. while Feature 54 produced a date of 1290 \pm 40 B.P. Dering (Appendix C, this volume) concludes that the few species of plants identified in the macrobotanical samples are consistent with the Tamaulipan Thornscrub community that has been present in the regions for the past 3,000 years. None of the samples that he analyzed came from the deepest recovery contexts and therefore no evidence corroborating the significant vegetation shift noted by the pollen/phytolith analysis exists.

Other Proxy Data Types and Trends

Molluscs

Seventy-one heavy fraction samples derived from the flotation of feature fill matrix were sorted for macrobotanical remains. It was hoped that these samples will produce a large collection of invertebrate and vertebrate fauna that could be employed as proxy data for interpretations regarding subsistence patterns and paleoclimate. To define some analytical units that would be meaningful in searching for diachronic trends in the data, we plotted the two sigma date ranges of each of the features. Based on this plot, we identified four clusters of date ranges with the oldest predating 2700 B.P. containing a single feature (Feature 7). Given the small number of faunal elements in the feature it was not used in subsequent analyses. The next oldest cluster dating from 2300-2700 B.P. included 15 features (11, 12, 74, 84, 121, 148-152, 154-156, 159, and 161). The youngest groups included six features (2, 16, 55, 57, 89, and 114) and post-dated 1700 B.P. The cluster of features dating between 1700-2300 B.P. contained 3 features (41, 75, 153) and faunal materials recovered from OZ B and OZ D, since organics from these two occupation zones were dated to 1900 \pm 40 B.P. (OZ D) and 1840 \pm 40 B.P. and 2010 \pm 40 B.P. (OZ B), respectively.

The samples of faunal remains derived from the features grouped into these three main clusters of units indicated that only seven species of invertebrate taxa occur in sample sizes equal to or greater than 20 minimum number of individuals (MNI). They are the Crossbar Venus (*Chione cancellata*), Bay Scallop (*Argopecten irradians*), Lightning Whelk (*Busycon sinistrum*), Oyster (*Crassostrea virginica*), Sunray Venus (*Macrocallista nimbosa*), and Quahog (*Mercenaria campechiensis*). *Rabdotus* (*Rabdotus* sp.) snails are also included in the figure given their correlation with climatic conditions (Brown 1999; Malof 2001).

In general, oysters, scallops, sunray venus, quahog and lightning whelk are considered moderate -to high-salinity species (see Table 8-5) preferring salinity ranges between 25-40‰. *Rabdotus* snails are commonly assumed to prefer semi-arid open grassy areas although they can also be found clinging to the stems of brushy vegetation such as found in the South Texas brush county (Personal Observations, Spring 2007, Wilson County).

Figure 10-3 is a plot of the MNI counts by species by time period. Several trends can be noted in this figure. For instance, oysters and Lightning Whelks decrease between 2300 and 1700 B.P. but increase after 1700 B.P. Sunray Venus and Quahog decrease over time, with the later decreasing dramatically between 2300 and 1700 B.P. Bay Scallops and Crossbar Venus fluctuate over time with their highest numbers being found between 2300-1700 B.P. *Rabdotus* snails decrease dramatically between 2300 and 1700 B.P. and continue to decline in numbers even after 1700 B.P.

What do these trends mean? We are aware that interpreting these patterns strictly from a paleoenvironmental perspective under estimates the possible effect of changes in hunter-gatherer exploitation practices upon the invertebrate fauna of the site. At the broadest level the dramatic decrease in oysters between the 2300-2700 B.P. and 1700-2300 B.P. temporal clusters suggests that conditions in Corpus Christi Bay may be changing from moderate salinity to significantly reduced salinity. This would be expected given the assumption that the barrier islands would have begun to form around 3000 B.P. and should have begun to decrease the salinity of the bay waters through reduced exchange with the open Gulf. The frequency of lightning whelk molluscs parallels the pattern in oysters, as does the occurrence of sunray venus between the two temporal clusters. Interestingly, the subsequent increase in oysters and lightning whelks after 1700 B.P. suggests that habitat conditions in the bays may have become more saline and closer to the preferred conditions of these species. Since the coalescence of the barrier islands should lead to decreased salinity within the bays, the sudden increase in the frequency of these species may reflect either a shift in the

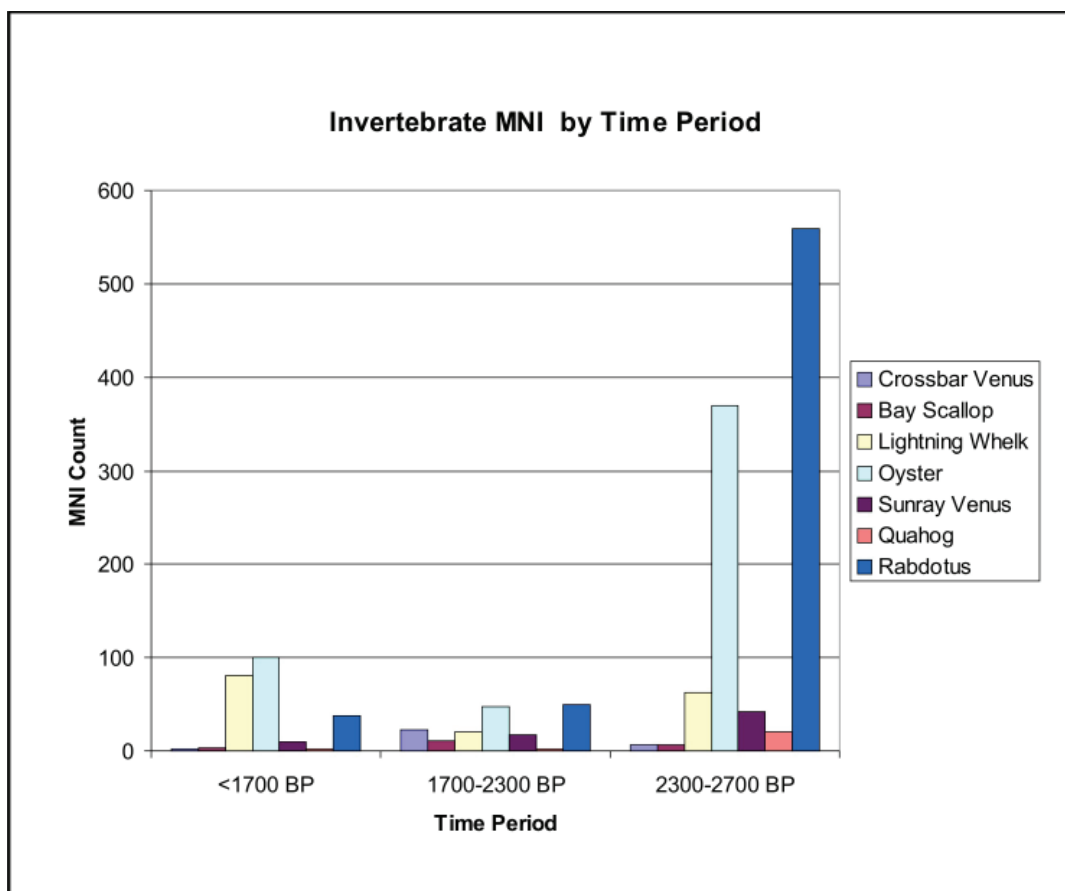


Figure 10-3. Plot of mni counts of seven invertebrate species by time period.

species targeted by foragers or perhaps a period of extended decrease in freshwater influx into the bays. It is worth noting that the most recent temporal cluster, is a composite of trends and therefore represents an average of a number of possible fluctuations in climatic conditions. With this in mind, we note that Core 2 and Core 6 particle size data obtained by both Frederick and Palacios-Fest seem to suggest that a period of steady fine-grained particle size deposition occurred beginning around 1770 \pm 40 B.P. and continued until around 1280 \pm 40 to 1300 \pm 40 B.P. Frederick's stable carbon isotope data also shows that until around 1000 B.P. the vegetation structure of the region was steadily moving toward an arid-adapted composition. Palacios-Fest's paleosalinity data from Core 6 also indicates that sometime around 1770 \pm 40 B.P. and until sometime before 1280 \pm 40 B.P. the salinity of the deposits was more marine than brackish. Finally, Cummings' et al. phytolith based climatic reconstructions also suggest that precipitation was decreasing around 1280 \pm 40 B.P. although it was apparently preceded by a period of increased precipitation between 2670 \pm 40 and 1770 \pm 40 B.P. In short, some degree of correspondence exists between the different data types in support of an extended period of decreased precipitation leading to increased bay salinity for some period

after 1770 \pm 40 B.P. Given this conclusion, the decrease in Rabdotos numbers over time is somewhat unexpected for two reasons. First, increased aridity and the establishment of more grass-dominated habitats should favor Rabdotos colonization. Second, the repeated reoccupation of the site by hunter-gatherers and the formation of several organically enriched paleosols and features should also have made conditions more rather than less favorable for the species over time.

An additional perspective on paleoclimatic conditions comes from a look at changes in species richness over time. As mentioned, the samples described above derive from a total of 26 features. Overall, the species richness among invertebrates within these 26 features fluctuates over time and this fluctuation does not seem to be related to sample size at least in terms of the number of features contributing specimens to the trend. For instance, the fifteen features grouped within the 2700-2300 B.P. temporal cluster have a combined richness of 26 species. The five features and two occupation zones grouped within the 2300-1700 cluster have a richness of 27 species. Finally, the six features post-dating 1700 B.P. have a richness of 19 species. Therefore, when considering all invertebrates present in these temporally specific subsamples,

it is clear that there is a significant decrease in species richness following 1700 B.P. This decrease may signal a reduction in the habitat conditions present within the bays exploited by the hunter-gatherers living at 41NU2. Such narrowing would result in the abandonment of the habitat by some species of invertebrates over time. Twelve species present in the samples between 2300 and 1700 B.P. drop out of the sample following 1700 B.P. Two of the new species (*Brachidontes* sp. and *Suturoglypta iontha*) that appear after 1700 B.P. were not present in the collections prior to 1700 B.P. The paleoclimatic implications of this trend are unclear at this time. One other aspect related to molluscs is worth mentioning, the entire absence of *Rangia* sp. molluscs from the 41NU2 sample. *Rangia* molluscs are brackish habitat preferring species and their absence is surprising given the site's proximity to what we assume to have been a brackish water bay. *Rangia* clamshells were common at the Mustang Lake site (41CL3) in Late Prehistoric deposits (Ricklis 2004:167) and their absence in an even more enclosed setting at 41NU2 suggests and implies that the salinity of the bays remained high throughout the sequence of human occupation especially considering how near the shore of Corpus Christi Bay the site is located and how small the freshwater contribution of the small ephemeral drainage at the south of the site may have been. Perhaps sites found closer to the mouth of the Nueces River would contain sizable *Rangia* samples.

Otholiths and Vertebrates

Unfortunately, while we hoped for large samples of otholiths from the heavy fraction samples, only 32 fish otholiths were identified during the sort of these samples. Fourteen (44%) of these were from unidentified bony fish leaving only 18 that were identified to taxa. The eighteen include Sea Catfish (n=7), Black Drum (n=3), Red Drum (n=2), Spotted Seatrout (n=2), Sand Seatrout (n=2), Sheepshead (n=1), and Atlantic Croaker (n=1). As the sample sizes show, none of these occurred in sufficiently large numbers to provide any meaningful statements regarding changes in their frequencies over time. Similarly, only 20 vertebrate faunal remains were identified to taxa. Again, none of the taxon occurs in sufficient numbers to describe statistically significant patterns.

Disagreements in Pollen Analysis

Finally, having discussed the various data types that contribute the picture on paleoclimate we have to also emphasize that major disagreements exist between the pollen data of Cummings et al. (Appendix A, this volume) and Vaughn Bryant's analysis (Appendix B, this volume). Specifically, Bryant's analyses concluded that the pollen samples from Core 2 represented highly degraded pollen assemblages that were not representative of the original vegetation communities that

inhabited the site and its environs. Therefore, he concluded that an extensive analysis of the samples was not warranted because it would result in a faulty and biased reconstruction. In addition, he suggested that the interpretation of highly degraded pollen samples is also fraught with danger in that pollen grains can be easily misidentified leading to further errors in reconstructions.

While Cummings et al. (Appendix A, this volume) have mentioned in several instances the degraded nature of the pollen samples particularly in sediment samples from Cores 4 and 6, they present a detailed reconstruction of paleoenvironmental conditions much of it based on the sequence from Core 2, and the specific samples that were judged too heavily degraded to warrant analysis. The differences in assessments, results and conclusions between the two laboratories are dramatic and necessitate systematic appraisal.

Summary

This chapter summarized the results of specialized analyses reported in several appendixes found at the back of this report including pollen and phytoliths (Appendix A and B), macrobotanical (Appendix C), diatoms (Appendix D), ostracods (Appendix E), and geoarcheological (Appendix F). We also attempted to merge the various results into one more or less corresponding reconstruction of paleoenvironmental conditions that existed at and in the vicinity of the site between 2700±40 and the present. Finally, we incorporated what evidence was available from other proxy data types that had implications for paleoclimatic reconstructions including primarily the molluscs. In the process we have provided a paleoclimatic record for the site and discussed the manner in which the apparent changes in conditions may have affected the available resources and/or themselves reflect the changes noted. The archeological record is simply not long enough to have provided information on sea level changes but it did implicate changes in near-site habitat conditions that affected the resource base particularly when it comes to the salinity of the bays and the molluscs populations they maintained. The stratigraphy of the site is composed of alternating organic-rich (darker colored) and organic-poor (lighter colored) zones. The darker zones clearly represent periods of dune stabilization and are the products of both anthropogenic as well as climatic factors.

The paleoclimatic record as it could be reconstructed from the composite of data types (which are not always in agreement), seems to show that early in the archeological sequence, perhaps around 4000 B.P., the site and its vicinity were dominated by more mesic-adapted plant communities but the

composition of this community began to shift to a greater proportion of xeric-adapted species sometime around 3000 B.P. This trend in composition, that also affected the structure of the communities, continued until about 1000 B.P. at which point there seems to have been a small shift back to mesic conditions. Albert's (2006) reconstruction of the period from 3050-0 B.P. characterizes it as shifting from mesic to sub-tropical. As such, this general reconstruction seems to agree with that seen in the 41NU2 record with the exception of the minor shift very late in the sequence.

The two pollen analyses conducted by Cummings et al. (Appendix A) and Bryant (Appendix B) come to dramatically different conclusions about the pollen assemblages present in the sites's sediments and specifically those from Core 2 recovered from the pond near the north end of 41NU2. The principal disagreements lie in the assessments of the

preservation conditions of the pollen assemblages and their feasibility for paleoenvironmental reconstructions. The disagreements severely undermine the results of the analyses and some of the main lines of evidence used in the reconstruction of paleovegetation at and near the site. The principal differences seem to derive from disagreements as to what level of pollen degradation is needed before an assemblage is considered unreliable in providing a representative reconstruction of past vegetation communities and paleoclimate. Both pollen extraction and identification methods also are implicated in this question. Therefore, given the importance of pollen studies to paleoclimatic reconstructions and the lack of reliable records in Texas outside of a handful of bogs, it is recommended that TxDOT-ENV pursue a systematic resolution of these disagreements to establish the source of the different study results and reestablish confidence in the methods employed in the extraction and analysis of pollen assemblages.

Chapter 11: 41NU2 and Regional Subsistence Practices

Barbara M. Meissner and Jennifer A. Thompson

Analysis of Faunal Remains from 41NU2

The cultural materials that were and are at the Cayo del Oso site consist overwhelmingly of faunal remains. Most features include large numbers of shell and/or bone. Some levels in the trenches were completely paved with layers of invertebrate shells.

The nature of the project at 41NU2, however, resulted in very limited collection of faunal materials. With the exception of the first 16 features, faunal remains were rarely collected except within soil samples. Even this collection has been limited, as only soil samples associated with features were floated, and the faunal materials in the heavy fraction analyzed. The fauna described in this section of the report, therefore, can be considered no more than an incomplete record of the fauna exploited by the early inhabitants of Cayo del Oso.

The taxonomic nomenclature used in this report is that of the Integrated Taxonomic Information System (ITIS). It is available online (ITIS 2008). In most cases, the names of taxa listed here are the same as those found in any faunal report of the coastal area. However, the taxonomic names of some species, in particular invertebrate species, vary a great deal in the literature. A good example is the Lightning Whelk, one of the most common shells found along the central Texas Coast. At various times the species has been referred to as *Busycon perversum*, *B. contrarium*, and *B. sinistrum*, all of which make reference to the fact that the Lightning Whelk is one of the few snail species in the world that twists to the left instead of the right. The confusion of names is the result of a certain amount of uncertainty on the part of biologists concerning how many species of Lightning Whelks there are, and where they are, but also because various analysts of faunal materials from coastal sites have used various reference materials as authorities. We will use *Busycon sinistrum* in this report, following ITIS (2008), though most of the interim reports used *B. perversum*.

Invertebrates

The vast majority of faunal remains collected during this project were the shells of mollusks. As noted in Chapter 2 and in other sections of this report, the estuarine environment in the immediate vicinity of 41NU2 provides an abundance of marine resources. The inhabitants of 41NU2 made use of many species, naturally concentrating on the larger and

most easily acquired. The floatation of a total 318.8 liters of soil samples from features resulted in the recovery of 16,665 identifiable invertebrates, weighing 26,935.38 g. In addition, a total of 1,114.35 g of shell could not be identified. Table 11-1 is a list of the 45 invertebrate taxa identified from the heavy fraction of soil samples, with counts and weights of each. The majority of the invertebrates were from the Phylum Mollusca (n=16,421), however there are also Crustaceans from the Phylum Arthropod (n=231), and Annelid worms from the Phylum Annelida (n=3) in this collection.

Table 11-1. Invertebrate Species Identified from Heavy Fraction Floatation Samples

Taxon	Common Name	Count	Weight (g)
<i>Phylum Mollusca</i>			
Class Bivalvia: Marine Bivalves			
<i>Abra aequalis</i>	Atlantic Abra	1	0.03
<i>Amygdalum papyrium</i>	Atlantic papermussel	90	6.71
<i>Anadara brasiliana</i>	Incongruous Ark	1	0.09
<i>Anadara transversa</i>	Transverse Ark	4	0.98
<i>Anodontia alba</i>	Buttercup Lucina	2	0.34
<i>Anomalocardia auberiana</i>	Pointed Venus	1	0.05
<i>Anomia simplex</i>	Common Jingle Shell	1	0.54
<i>Argopecten irradians</i>	Atlantic Bay Scallop	252	148.52
<i>Brachidontes sp.</i>	Scorched Mussels	3	0.66
<i>Carditamera floridana</i>	Broad-Ribbed Cardita	1	0.05
<i>Chione cancellata</i>	Crossbar Venus	169	103.96
<i>Codakia oribicularis</i>	Tiger Lucine	1	0.03
<i>Crassostrea virginica</i>	Eastern Oyster	8,222	11,101.58
<i>Dinocardium robustum</i>	Giant Atlantic Cockle	74	58.60
<i>Diplodonta punctata</i>	Atlantic Diplodon	1	0.25
<i>Dosinia discus</i>	Disk Dosina	1	0.03
<i>Eucrassatella speciosa</i>	Beautiful Crassatella	1	0.04
<i>Ischadium recurvum</i>	Hooked Mussel	42	3.39
<i>Lucine amianta</i>	Lovely Lucine	2	0.05
<i>Macoma constricta</i>	Constricted Macoma	2	0.02
<i>Macrocallista nimbosa</i>	Sunray Venus	2,614	4875.03
<i>Mercenaria campechiensis</i>	Southern Quahog	344	1357.34
<i>Mulina lateralis</i>	Dwarf Surfclam	1	0.08
<i>Mytilidae</i>	Saltwater Mussels	52	34.43
<i>Noetia ponderosa</i>	Poderous Ark	2	0.61
<i>Nuculana acuta</i>	Pointed Nutclam	2	0.02
<i>Semele bellastrata</i>	Cancellate Semele	2	0.02
<i>Spisula raveneli</i>	Southern Surfclam	1	0.02
<i>Strigilla carnaria</i>	Large Strigilla	1	0.78
<i>Trachycardium sp.</i>	Prickly Cockles	4	0.78
<i>Tracycardium egmontianum</i>	Florida Prickly Cockles	1	0.66
Total Marine Bivalves		11,895	17,695.69
Class Gastropoda: Marine Gastropods			
<i>Busycon sinistrum</i>	Lightning Whelk	1,638	7451.46
<i>Busycotypus canaliculatus</i>	Channeled Whelk	5	59.56
<i>Busycotypus spiratus</i>	Pear Whelk	4	2.91

Table 11-1. Continued...

Taxon	Common Name	Count	Weight (g)
Phylum Mollusca			
Class Gastropoda: Marine Gastropods			
<i>Busycon sinistrum</i>	Lightning Whelk	1,638	7451.46
<i>Busycotypus canaliculatus</i>	Channeled Whelk	5	59.56
<i>Busycotypus spiratus</i>	Pear Whelk	4	2.91
<i>Cerithiopsis greenii</i>	Green's Cerith	1	0.03
<i>Crepidula depressa</i>	Faded Slippershell	21	2.16
<i>Crepidula fornicata</i>	Common Slippershell	34	1.82
<i>Crepidula plana</i>	Eastern White Slippershell	1	0.03
<i>Crepidula</i> sp.	Slippershells	1	0.02
<i>Fasciolaria lilium</i>	Banded Tulip	38	127.79
<i>Marshallora nigrocinta</i>	Black Lined Triphora	2	0.03
<i>Pleuroploca gigantea</i>	Florida Horse Conch	12	93.26
<i>Polinices duplicatus</i>	Shark's eye	30	159.52
<i>Suturoglypta iontha</i>	Lineate Dovesnail	3	0.1
<i>Terebra dislocata</i>	Common Atlantic Auger	2	0.21
Total Marine Gastropods		1,792	7,898.90
Class Gastropoda: Terrestrial Gastropods			
<i>Rabdotus</i> sp.	Rabdotus snails	2,733	1315.32
Total Gastropods		4,525	9,214.22
Class Scaphopoda			
<i>Dentalium americanum</i>	American Tusk shell	1	0.03
Total Mollusks		16,421	26,909.94
Phylum Arthropoda			
Balanomorpha	Acorn Barnacles	229	0.22
Decopoda	Crabs	2	24.13
Total Arthropods		231	24.35
Phylum Annelida			
Class Polychaeta: Polychete worms			
<i>Sabellaria</i> sp.	Sandtube Worms	3	1.09
Total Annelids		3	1.09
Total Invertebrates		16,655	26,935.38

The most common shells recovered during processing of the heavy fractions were Eastern Oysters (*Crassostrea virginica*; n=8,222), Sunray Venus (*Macrocallista nimbosa*; n=2,614); *Rabdotus* land snails (n=2733) and Lightning Whelks (*Busycon sinistrum*; n=1,638). Together these four species make up 91.3 percent of the total invertebrates in this collection and 91.9 percent of the identified shell weight. With the exception of the *Rabdotus*, this is not surprising, as these are all quite large species, relatively easy to acquire. Other species large enough to eat were collected and brought to 41NU2 as well. As Table 11-1 shows, Atlantic Bay Scallops (*Argopectin irradians*), and the Southern Quahog (*Mercenaria campechiensis*) were exploited. Another, less common, large species was the Giant Atlantic Cockle (*Dinocardium robustus*). The Florida Horse Conch (*Pleuroploca gigantea*), though by far the largest invertebrate available to the inhabitants of 41NU2 (Andrews 1977: 146), is not common in the Corpus Christi Bay area today and the relatively small numbers (n=12) recovered from the soil samples suggests that this was true in the past as well. On the other hand, the Sunray Venus clam, the second most-commonly identified shell in the collection, no longer

lives along the Texas coast (Andrews 1977:242), though it is still well represented in the waters of Florida's bays and a new industry of commercial farming of the species is under development there (Sturmer 2007).

The presence of large number of *Rabdotus* shell in some of the features is not easy to explain. These land snails are much smaller than most of the marine species in terms of the amount of food-value per individual. While they could represent low ranked food items (Malof 2001) we did not conduct research to attempt to test this possibility associated with this project. The fact that these species feed on and are attracted to organic-rich deposits may also indicate that the clusters of snails in features may represent die-offs rather than discarded consumption debris.

Soils matrix from 61 features were floated and Table 11-2 shows the number of samples in which each taxon was identified, and the percent of the 61 heavy fraction samples in which each was noted. If we look at the relative ubiquity of the 45 species recovered, we can see that, again, oysters were the most ubiquitous species. The presence of acorn barnacles in 18 of the samples and in relatively large numbers (n=229) is interesting. They were probably attached to the oysters that were collected, and may even have been eaten along with other invertebrates collected during trips to the oyster reefs.

Another species that is fairly ubiquitous is the Crossbar Venus. This relatively small clam is often no bigger than 15 mm and can probably serve as a lower size boundary of species that might have been actively gathered for food, though the low number of shells in most features suggests they were usually picked up incidentally, as a result of their association with oyster reefs (Kilgen and Dugas 1989:45).

Twenty-two species were found only in a single soil sample. Of these, it is interesting that 8 were found only in Feature 41. In fact, Feature 41 had 24 identified species, with a total count of 1,022 identified shells. Feature 41 was a small pit, about 30 cm on a side, and only about 5 cm deep (see Figure 6-19). The remarkable number and diversity of species found in this one feature make it unique.

Vertebrates

Although at least some bone was visible in most features and in large areas of Occupation Zones, it was for the most part fragmented into tiny pieces, and much of it was the bone of small fish that more or less disintegrated when disturbed, no matter how carefully the bone was handled. Even bones of large, deer-sized mammals were usually highly fragmented

Table 11-2. Invertebrate Taxa from 41NU2, Showing the Ubiquity of Each, and Noting Association with Oyster Reefs

Taxon	Number of Features Where Present	Percent of Features	Found on Oyster Reefs
<i>Crassostrea virginica</i>	45	73.8%	Yes
<i>Busycon sinistrum</i>	44	72.1%	Yes
<i>Rabdotus sp.</i>	38	62.3%	
<i>Macrocallista nimbosa</i>	36	59.0%	
<i>Balanomorpha</i>	18	29.5%	Yes
<i>Mercenaria campechiensis</i>	17	27.9%	Yes
<i>Argopectin irradians</i>	11	18.0%	
<i>Polinices duplicatus</i>	11	18.0%	Yes
<i>Chione cancellata</i>	10	16.4%	Yes
<i>Dinocardium robustum</i>	10	16.4%	
<i>Crepidula fornicata</i>	8	13.1%	Yes
<i>Decapoda</i>	8	13.1%	
<i>Amygdala papyrium</i>	7	11.5%	
<i>Faciolaria lillium</i>	7	11.5%	
<i>Pleuroplaca gigantea</i>	7	11.5%	
<i>Crepidula depressa</i>	5	8.2%	Yes
<i>Mytilidae</i>	5	8.2%	Yes
<i>Ishadium recurvum</i>	4	6.6%	Yes
<i>Sabellaria sp.</i>	3	4.9%	Yes
<i>Anadara transversa</i>	2	3.3%	Yes
<i>Brachidontes sp.</i>	2	3.3%	Yes
<i>Marshallora nigracintus</i>	2	3.3%	
<i>Trachycardium sp.</i>	2	3.3%	Yes
<i>Abra aequalis</i>	1	1.6%	Yes
<i>Anodontia alba</i>	1	1.6%	
<i>Anomalocardia auberiana</i>	1	1.6%	
<i>Anomia simplex</i>	1	1.6%	Yes
<i>Busycotypus canaliculatus</i>	1	1.6%	
<i>Busycotypus spiratus</i>	1	1.6%	
<i>Carditamera floridana</i>	1	1.6%	
<i>Codakia oricularis</i>	1	1.6%	
<i>Dentalium americanum</i>	1	1.6%	
<i>Diplodonta punctata</i>	1	1.6%	
<i>Dosinia discus</i>	1	1.6%	
<i>Eucrassatella speciosa</i>	1	1.6%	
<i>Lucine amianta</i>	1	1.6%	
<i>Macoma constricta</i>	1	1.6%	
<i>Mulina lateralis</i>	1	1.6%	Yes
<i>Noetia ponderosa</i>	1	1.6%	Yes
<i>Nuculana acuta</i>	1	1.6%	
<i>Semele bellastrata</i>	1	1.6%	
<i>Spisula raveneli</i>	1	1.6%	
<i>Strigilla carnaria</i>	1	1.6%	
<i>Suturoglypta iontha</i>	1	1.6%	
<i>Terebra dislocata</i>	1	1.6%	

(see Figure 6-23). Yet the few large pieces of bone from mammals were often found in good to excellent condition, suggesting that the mammal bone had been fragmented as part of the butchering process and before deposition.

The total of identified bone from the soil samples that were floated was 3,717 pieces amounting to a total identified bone weight of 362.9 g. Table 11-3 shows a complete list of vertebrate remains that could be identified in feature soil samples. The vast majority of bone from the heavy fraction was in pieces less than 20 mm and was often hardly more than bone “dust”. No attempt was made to count unidentifiable bone, though it was weighed and totaled 241.70 g.

Table 11-3. Vertebrate Species Identified from Heavy Fraction of Floatation Samples

Taxon	Common Name	Count	Wgt. (g)
Mammals			
Bovinae	Bison or Cattle	3	7.3
<i>Odocoileus virginianus</i>	Whitetailed Deer	12	64.62
Artiodactyla	Deer or antelope-sized	8	40.85
<i>Taxidea taxus</i>	American Badger	1	0.07
Carnivora	Unidentified Carnivore tooth	1	0.05
<i>Neotoma sp.</i>	Packrats	3	0.12
<i>Sigmodon hispidus</i>	Hispid Cotton Rat	1	0.03
Rodentia	Unidentified Rodents	21	0.54
Mammal-V. Lg.	Bison, Cattle, or Horse-sized	30	18.12
Mammal-Lg.	Deer-sized	42	38.74
Mammal-Medium	Dog, Coyote-sized	5	29.1
Mammal-Sm.	Rabbit-sized	1	1.03
Mammal-V. Sm.	Rat-sized	2	0.05
Mammal	Size Indeterminate	15	0.58
Total Mammals		145	201.2
Birds			
Aves-Small	Robin-sized	2	0.06
Aves	Size Indeterminate	26	0.78
Total Birds		28	0.84
Reptiles			
Serpentes	Unidentified Snake	1	0.02
Testudines	Unidentified Turtle	5	0.58
Total Reptiles		6	0.6
Fish			
<i>Archosargus probatocephalus</i>	Sheepshead	2	0.34
<i>Ariopsis felis</i>	Sea Catfish	27	4.04
<i>Bairdiella chrysoura</i>	Silver Perch	1	0.06
<i>Cynoscion arenarius</i>	Sand Sea Trout	14	1.51
<i>Cynoscion nebulosa</i>	Spotted Sea Trout	3	0.23
<i>Micropogonias undulatus</i>	Atlantic Croaker	2	0.3
<i>Pogonias cromis</i>	Black Drum	5	3.29
<i>Sciaenops ocellata</i>	Red Drum	7	0.75
Osteichthyes	Unidentified Boney Fish	3,477	149.74
Total Fish		3,538	160.26
Overall Total		3,717	362.90

In the field it was possible to see in some features layers of fish bone as much as 50 mm thick spread across areas of a meter or more. The diameters of the centric of the vertebrae in these lenses of fish bone usually were no more than 30 mm suggesting fairly small fish species. The presence of such dense layers of recognizable fish bones does suggest that the fish has undergone relatively little processing before consumption. Reduced processing costs, within the context of mass-procurement would make such resources relatively highly ranked as mentioned before. Unfortunately, while these dense deposits of fish bones represent significant information, the data becomes what Lyman and O'Brien (1987) called "analytically absent" because all attempts to remove them from the surrounding clay matrix or otherwise quantify them based only a profile exposure either fail or are inadequate. Even attempts to photograph these beds of fish bone were not very successful. It is not possible, therefore, to quantify, even very roughly, the degree to which the inhabitants of 41NU2 made use of small fish in their diet. Fish comprise 95.2 percent of the identifiable vertebrate bone (Table 11-3), and given all data available, both quantitative and subjective observation in the field, it is likely that, at least at times, fish was an even more important part of the diet than shell fish.

Discussion

In considering the result of the analysis of the faunal material in the heavy fractions of the soil samples from 41NU2, the limitations of the data, especially concerning vertebrate fauna, are striking. The soil samples were often only a very small percentage of the total feature, and since larger features were not always uniform in structure or contents, the samples do not always reflect the total species diversity and relative dietary importance. This is especially true in features where there were a great number of small fish bones, which could not be adequately enumerated. Oso Bay would have been a perfect place for netting, using weirs, or even scooping into baskets to collect the many species of fish that spend their juvenile years in these shallows, but we have only the massive numbers of tiny fish bones observed in the field, which cannot be enumerated in a satisfactory way, to show that this was done.

The fish species identified at 41NU2 area listed in Table 11-4, with information on the annual presence/absence of these species in the bays. In six features (Features 73, 75, 114, 117, 155, and 156), two or more species of fish were found together. Comparing the presence data for the fish in question (Table 11-4) allows us to suggest when these features might have been in use (Table 11-5). As Table 11-5 shows, all of these features are likely to have been utilized during the spring. This data is insufficient to suggest with any degree

Table 11-4. Seasonal Presence and Absence in Estuaries for Eight Fish Species

Species	Presence/Absence in Bays
Sheepshead	Absent mid-summer into fall (Jun.-Nov.)
Sea Catfish	Absent mid-summer into fall (Jun.-Nov.)
Silver Perch	Year round, into shallows and rivers in summer
Sand Sea Trout	Absent only in mid-winter (Nov.-Feb.)
Spotted Sea Trout	Juveniles remain year-round; Adults move in and out depending on temperature and salinity
Atlantic Croaker	Late spring through summer
Black Drum	Juveniles all year; Adults absent Feb-April
Red Drum	Juveniles all year; Adults absent July-Jan.

Table 11-5. Six Features, the Multiple Fish Species Present, and Estimate Months of Utilization

Feature	Fish Species Present	Estimated Months of Utilization
73	Sea Catfish; Red Drum	Feb. to April
75	Sheepshead; Sea Catfish; Spotted Sea Trout	Dec. to April
114	Sea Catfish; Spotted Sea Trout; Red Drum	Feb. to April
117	Sea Catfish; Silver Perch; Sand Sea Trout Spotted Sea Trout	March to April
155	Sea Catfish; Spotted Sea Trout; Red Drum	March to April
156	Spotted Sea Trout; Red Drum	March to April

of certainty when in the annual round 41NU2 was occupied or if the season of occupation changed over time. Several other forms of seasonality estimation and more dating of features would be required to answer this question with any confidence. However, two other types of evidence do suggest that the site was frequently inhabited during a time of the year when large-bodied food resources were scarce or absent, forcing the inhabitants to turn to procurement of resources that would be far lower ranked.

The first indication that habitation at 41NU2 was frequently during the late winter and/or early spring is the presence of the large numbers of tiny fish bones evident in many of the features. As has been noted above, these remains of these fish were observed frequently in the field, often in a thin layer of extremely dense fishbone, all or most of which were very small, with most representing fish less than 10 cm in length and many at least half that size. Clearly, catching such fish on an individual basis would be energetically far too expensive and would yield low returns. However, the application of technology, in the form of basket scoops can increase the capture rate significantly, although it may also increasing overall technology costs.

(Bettinger et al. 2006; Ugan 2005:80; Ugan et al. 2003). The fact that such small fish were probably eaten whole, with or without cooking, thus reducing processing cost significantly, adds to the potential return rate (Ugan 2005:86). The presence of so many small fish in multiple features across the site also may suggest that other, more highly ranked food resources were not available or were scarce.

Another indication that habitation at 41NU2 often took place during times of the year when many food resources were scarce is the fact that in most features and Occupation Zones, deer-sized animal bone is present, but has been reduced to small splinters. The undisturbed, well-sealed context in which most features were found, and the many small hearths and small shell clusters indicate that any given habitation at 41NU2 was usually brief. It is unlikely, then, that the extreme fragmentation of most animal bone seen on the site was caused by trampling or by post-depositional disturbances. It appears, therefore that the bone was fragmented as part of the butchering process. While not enumerated, field observations and examination of the small percentage of bone that was returned to the lab and analyzed supports the hypothesis that most of the fragmentation was done while the bone was fresh. This may be an indication that those portions of large-bodied animals that were brought to the site were processed for bone grease.

The extraction of bone grease involves breaking fat-rich elements, such as the ends of long bones, into small pieces and then boiling the bone fragments until the fat in them melts and floats to the surface of the water, where it is scooped up (Brink 1997; Vehik 1977). This is a labor and fuel intensive process with a relatively small yield. Yet Brink (1997:271) has pointed out that at certain times of the year in particular, the need for even a little fat can become paramount. The human body cannot process protein without fat (Brink 1997; Speth and Spielman 1983:149-159). Since most animals are very lean during the late Winter/early Spring, or during prolonged droughts, it is during these times that the fat in bones will seem important enough to make the effort (Brink 1997:271). Vehik (1997) has presented characteristics one would expect to see if bone has been processed for bone grease. The characteristics include large numbers of very small fragments of bone which appear to be from deer or bison-sized animals and an absence or near absence of identifiable long bone ends and the centrums of vertebrae. Although the overall faunal collection at 41NU2 does contain some bison-sized bone found in relatively large pieces, the vast majority of the bones of deer or larger mammals are extremely fragmented and very few fragments of articular ends were observed.

Examining the small quantity of mammalian bone is of limited value, as the sample size is so small and since provenience information indicates it comes from several

thousand years of occupation. By grouping mammals by size class, that is by bison-sized, deer-sized and smaller than deer-sized, and by comparing the bone weights of these groups as a percentage of their total bone weight, we can state with some confidence that deer-sized animals are the most common fauna represented at 41NU2 (Table 11-6).

Table 11-6. Mammalian Bone Weight in Three Size Categories

Body-Size Category	Bone Wgt. (g)	Percent of Total
Bison-sized	25.42	14.8%
Deer-sized	144.21	84.1%
Smaller than Deer-sized	1.89	1.1%
Total	171.52	100.0%

Shell fishing seems to have been the most important gathering activity practiced by inhabitants, though the few features that consisted largely of bone suggest that this was not always the case. The inhabitants of 41NU2 had two major shell fishing options, the oyster reefs in Corpus Christi Bay, and the mud bottoms and sea grass beds of Oso Bay. In fact, of the species found in at least 10 soil samples, the only species not commonly associated with oyster reefs are large species with considerable food value, and the *Rabdotus* land snail (Table 11-2). The only major species that is not associated at least partially with oyster reefs is the *Sunray Venus*. This clam would have likely been common in the mudflats of Oso Bay. Lightning Whelks would have preyed on both oysters and on the bivalves in mudflat and sea grass beds, and would thus have been readily collected while looking for clams or while harvesting the oyster reefs. Although there is no oyster reef near Oso Bay in Corpus Christi Bay today, as recently as 1961 two oyster reefs were identified just offshore of the entrance to Oso Bay (Martinez 1961:2). Martinez (1961:1) noted that most of the oysters he encountered during his survey of Corpus Christi Bay were dead. At that time the only viable reef (i. e. a reef large enough to warrant commercial shellfishing) was at the northern end of the bay, near the entrance to Nueces Bay (Martinez 1961:1). Martinez (1961) attributed the small number of living oysters to the use of dredges for oyster collecting, which had been the practice for some 70 years (Townsend 1893:414) and had only recently been discontinued (Martinez 1961:1). Another possible cause may have been a recent large influx of freshwater from storms. The salinity of Corpus Christi Bay is often lowered to the point that oysters cannot survive after such an influx (Townsend 1893:414).

Comparative Sites

In order to compare the subsistence practices observed at 41NU2 with other sites in the region, as defined in the

Research Design (Tomka et al. 2006), the work authorization for this report required a detailed literature review and collection of data on 13 sites in the general vicinity of 41NU2: 41AS3, 41NU184, 41NU221, 41NU266, 41SP15, 41SP43, 41SP120, 41SP136, 41SP148, 41SP149, 41SP153, 41SP156, and 41SP177.

During the examination of the literature it was discovered that no report concerning an excavation at site 41SP15 was available, other than mention of a brief visit and acquisition of a sample for radiocarbon dating (Ricklis 1993). While attention to vertebrate fauna in many of the other sites was minimal, most provided data on the numbers of invertebrates at the site, though in most cases, only large and/or readily identifiable species were reported. In one case, site 41AS3, the available reports did not have an analysis of the faunal material, merely mentioning the presence of a given species (Campbell 1955; Cox and Smith 1988; Jurgens and Whitsett 1982; Ward and McGraw 1993). Since neither 41AS3 nor 41SP15 had useable faunal data, they were removed from this study.

In order to make the comparison of 41NU2 with other sites more comprehensive, site 41NU11, which had excellent data available (Hedrick 1993), was added. Site 41AS16, located on low mounds in a marshy grass flat area in Copano Bay was also added (Prewitt et al. 1984). In addition, the report on the excavation of 41NU266 showed that sites 41NU267 and 41NU268 were less than 100 m from 41NU266 and that 41NU278 was closer to Locus 3 in 41NU266 than Locus 1 and 2 were. In order to add to the sample sizes, these three sites, all occupying the same knoll as 41NU266, were also added to the database.

Thus, in the remaining portion of this chapter the faunal data from 41NU2 will be combined with sixteen other sites: seven in the Bay-Fluvial habitat, two in the marshy Grass Flats habitat, and seven in the Prairie-Riverine habitat. Table 11-7 shows the sites chosen for the comparative study, including a brief description of their location, the habitat category in which they have been placed, and the literature from which the data was acquired. Appendix H presents the complete faunal database from the 17 sites, including 41NU2, used in the following discussion. The data presented in Appendix H include: site number, time period, taxa, and counts either as NISP or MNI when reported. The mammals and birds were grouped by body size. Reptiles were reported by Order and Suborder (i.e. "Turtles", "Snakes", and "Lizards"). Invertebrates were reported by Genera or Species whenever possible.

The three site habitats mentioned above describe in a general way the setting in which the sites were located. Sites 41AS16 and 41SP177 were located on low mounds at the edge of a bay (Copano and Nueces, respectively), surrounded by extensive salt marshes, in the Grass Flats habitat. Sites 41NU11, 41NU184, 41NU221, 41NU266, 41NU267, 41NU268, and 41NU278 were located on terraces above the Nueces River floodplain. Though all are today some distance from the mouth of the river, the original valley was both deeper and narrower, and it was the development of the Nueces delta and the extension of that delta towards Nueces Bay that has resulted in the site settings seen today. In the past, the river mouth would have been closer and the water more brackish than today, as evidenced by the presence of oyster and Lightning Whelk, as well as the low salinity Rangia

Table 11-7. Sites Used in Comparison of Regional Subsistence Practices, Including Their Habitat and References Used for Data Collection

Site	Description of Location	Site Habitat	References
41AS16	On two relict Pleistocene knolls surrounded by low salt marsh at the point where Swan Lake enters Copano Bay;	Grass Flats	Prewitt et al. 1984; Paine 1987
41NU11	On a clay dune located just south of the entrance of Oso Creek into Oso Bay;	Prairie-Riverine	Headrick 1993
41NU184	Uplands of dissected Pleistocene terrace above floodplain of Nueces River;	Prairie-Riverine	Ricklis and Gunter 1984
41NU221	Uplands overlooking the modern Nueces River floodplain;	Prairie-Riverine	Ricklis 1986, 1988
41NU266/NU267/ NU268/NU276	Uplands overlooking the modern Nueces River floodplain. The 4 sites are all within 100 m of each other;	Prairie-Riverine	Ricklis 1995; Warren 1993, 1994
41SP43	On low bluff immediately adjacent to Ingleside Cove (Corpus Christi Bay). Sites SP43 and SP120 are probably the same occupations;	Bay-Fluvial	Corbin 1963; Ricklis 1993; Story 1968
41SP120	On low bluff above Ingleside cove (Corpus Christi Bay). Sites SP43 and SP120 are probably the same occupations;	Bay-Fluvial	Prewitt 1984; Ricklis 1990
41SP136	Uplands immediately adjacent to partially infilled cove of Nueces Bay;	Bay-Fluvial	Ricklis 1993
41SP148	Uplands immediately adjacent to partially infilled cove of Nueces Bay;	Bay-Fluvial	Ricklis 1993
41SP149	Uplands immediately adjacent to partially infilled cove of Nueces Bay;	Bay-Fluvial	Ricklis 1993
41SP153	Uplands immediately adjacent to partially infilled cove of Nueces Bay;	Bay-Fluvial	Ricklis 1993
41SP156	On knoll that is remaining portion of Pleistocene terrace;	Bay-Fluvial	Ricklis 1993
41SP177	Marshy area at foot of Pleistocene uplands, near small creek;	Grass Flats	Ricklis 1993

clams found, sometimes in great numbers, at these sites (Ricklis 1995, 1986, 1988, 1990; Ricklis and Gunter 1986). These Prairie-Riverine habitat sites had the resources of both the river and environs, and the prairies that surrounded them. Only in the earliest levels of these sites are more marine fauna seen. This will be considered below.

Sites 41SP43, 41SP120, 41SP136, 41SP148, 41SP1149, 41SP1153, and 41SP156, as well as 41NU2, are located on bluffs above a bay, in areas where there is a freshwater influence nearby. These Bay-Fluvial habitat sites can take advantage of the increased marine resources available in such habitats (Stickney 1984:39).

Analytic problems

It is important to understand that the level of reporting of faunal remains in the literature concerning these sites is extremely variable. This is largely due to three factors: a change over time in the understanding of the importance of faunal analysis in the investigation of past lifeways; the availability at the time of faunal experts to make the analysis; and the difficulties of finding funding for such analyses. As a result the level of expertise used in identification of the animal remains at the sites under consideration was variable. This is an unfortunate, but inevitable, consequence of using data from more than fifty years of reports, and from both relatively amateur to professional faunal analysts. Short of the reanalysis of the collections chosen for this comparative study by a single analyst or team of analysts using the same criteria, there is no way around this problem.

In addition it must be noted that using NISP as a measure of dietary importance has some obvious drawbacks. NISP is simply the number of identified specimens in a given sample. From NISP you cannot determine if the 20 bison bones came from 1 or 20 animals. More importantly, the NISP of 20 bison bones and 20 rat bones suggests a diet comprised of equal amounts of bison and rat. Clearly this is not the case, as each bison bone represents a much larger amount of meat than each rat bone.

Various ways to compensate for this analytical weakness of NISP have been devised. In certain limited circumstances, the Minimum Number of Individuals (MNI) can be a useful method for reconstructing the diet represented by bone recovered from archaeological sites. In most circumstances the use of bone weight is probably the best method to gauge the relative importance of various mammalian species to a diet. If large numbers of non-mammalian species are present, it is best to use the allometric formulae devised by Reitz (Reitz and Wing 1999:22-231) to achieve

the best fit of estimated biomass (and thus relative dietary importance). Unfortunately, very few of the reports listed in Table 11-7 mention the weight of the bone recorded. We are forced to use NISP, with, perhaps, an eye to the relative size of the animals in question. Therefore, all identified mammal bone was divided into five body-size classes, representing bison-sized, deer-sized, coyote-sized, rabbit-sized, and rat-sized animals.

The invertebrates recorded in the reports listed in Table 11-7 are different. In each report, shell fragments that did not contain a substantial portion of the umbo or columnella of the shell were not counted. Thus the counts are either the MNI in the case of Gastropods, or ½ the MNI in the case of Bivalves. It would probably be better to weigh the shells in this case, as well, to get a better sense of the overall importance of various species, as oysters that were no bigger than a thumb nail would be counted the same as oysters 20 cm long. An estimate of “meat-weight” represented by a given shell might be possible, if data on this can be located for enough of the most common species.

In any case, comparing the bones of vertebrates with the shells of invertebrates is not straightforward. In the circumstances, NISP is probably the best method to use, but it should not be supposed that if a given site has 100 mammal bones and 900 shells, that shellfish were nine-tenths of the diet. Instead, we can examine the differences, first by habitat, and then by time, in the relationship between the animal remains in the various vertebrate and invertebrate categories.

Comparing Subsistence in Three Habitats

Vertebrates

There are a number of ways to compare the evidence of subsistence identified in the sites listed above. To begin with, we can look at the vertebrate species. Table 11-8 shows the combined total of vertebrates, excluding fish, from the sites in each category for each habitat. Fish will be considered separately, as they are (with one exception, see below) an exclusively marine resource, while the other vertebrates are terrestrial or riverine resources. A possible exception to this would be marine turtles, however, no marine turtles were identified at any of the sites.

Figure 11-1 shows the difference in the relative use of three classes of vertebrates, mammals, birds and reptiles, in the site grouped into the three habitat types. The sites in the Prairie-Riverine habitat have the smallest total numbers, with Mammals (53.2 percent) and Reptiles (43.1 percent) dominating (Figure 11-1). The Grass Flats sites show a similar

Table 11-8. Counts of Selected Vertebrates from 17 Coastal Sites Grouped by Habitat

Selected Vertebrates*	Site Habitat		
	Grass Flats	Bay-Fluvial	Prairie-Riverine
V. Large (Bison-sized)	0	42	31
Large (Deer-sized)	115	766	30
Medium (Coyote-sized)	0	11	2
Small (Rabbit-sized)	10	13	9
Very Small (Rat-sized)	19	79	28
Total Mammals	144	911	100
V. Large (Turkey, Heron-sized)	0	0	1
Large (Duck-sized)	1	9	2
Small (Robin-Sized)	0	2	0
Total Birds	1	11	3
Alligator	1	1	2
Turtles	140	47	68
Snakes	137	21	15
Lizards	1	0	0
Total Reptiles	279	69	85
Total of Selected Vertebrates	424	991	188

*excludes fish

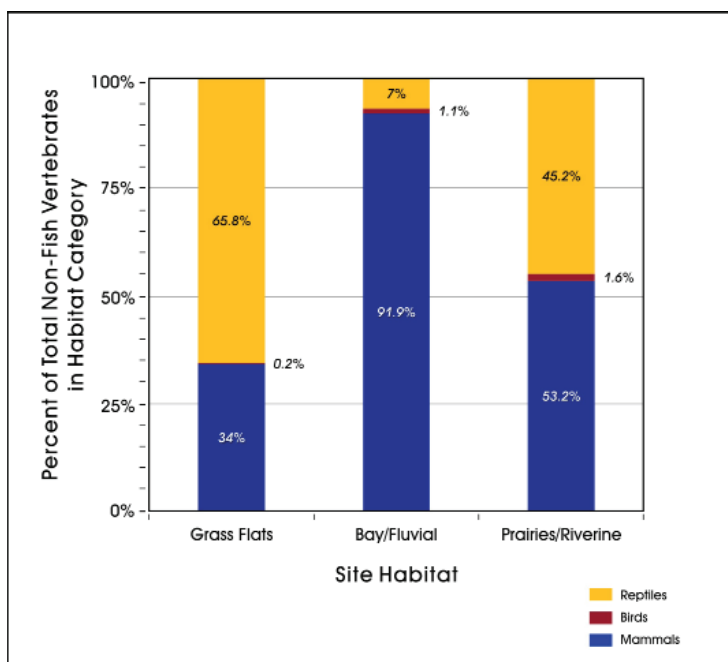


Figure 11-1. Percentage of identified vertebrate species (excluding fish) from 41NU2 and 17 comparative sites grouped into habitat type categories.

picture, with Mammals and Reptiles together comprising all but 0.2 percent of the total (Figure 11-1). The sites in the Bay-Fluvial areas are very different. In these sites 91.9 percent of the selected vertebrates are in the Mammal category, Reptiles are only 7.0 percent and birds are 1.1 percent.

The relatively low numbers of birds at all the sites was also noted at 41NU2. Though large numbers of birds are present in the area today, especially during nesting season and during the spring and fall migrations, very few bird bones were recovered from any of the sites in the comparative data set. At 41NU2, not only were there very few bird bones recovered (n=28), but no egg shell was found. As noted in Chapter 8, the numbers of birds, and the richness of species varies a great deal over the year, yet even in the period when the bird populations are smallest (i.e. April to July) bird species, many of which are in the medium to large category, are still available (Carrilo 2000; see also discussion in Chapter 8).

The question of why so few bird remains are found in these coastal sites may have to do with the difficulty of catching them, compared to the relative ease of digging clams and oysters out of the bays, and it may not be a coincidence that it is the Prairie-Riverine habitat, where mollusks are limited to the narrow river and thus are easily over-exploited, that bird bone was the most common (Figure 11-1). In this environment, mollusks of reasonable size would have been much less readily available than they are in either the Grass Flats or Bay-Fluvial habitats.

This might place birds slightly higher on the resource status list when these sites were inhabited.

A closer look at the kinds of Mammals being hunted in the three habitats reveals an interesting pattern. Figure 11-2 shows the breakdown of the mammal category by body-size. In this case, it is the Grass Flats and Bay-Fluvial habitats that are very similar in the percentage of each body-size, though the Bay-Fluvial category shows more presence of rabbit- and rat-sized bones than the Grass Flats Category. The Prairie-Riverine habitat bears no resemblance to the others. Relatively large percentages of bison- and deer-sized bones are present, along with a percentage of rat-sized bones that is nearly as large as the bison-sized and deer-sized bones. It is clear that close to the bays, the emphasis in mammal hunting is on deer-sized animals, with other mammals a more or less incidental presence. In the Prairie-Riverine sites, bones come from all categories, especially bison-, deer-, and rat-sizes. If we presume that bison would always be high on the resource status list, with deer and similar-sized animals considered secondary, whenever bison were available, Figure 11-2 suggests that bison

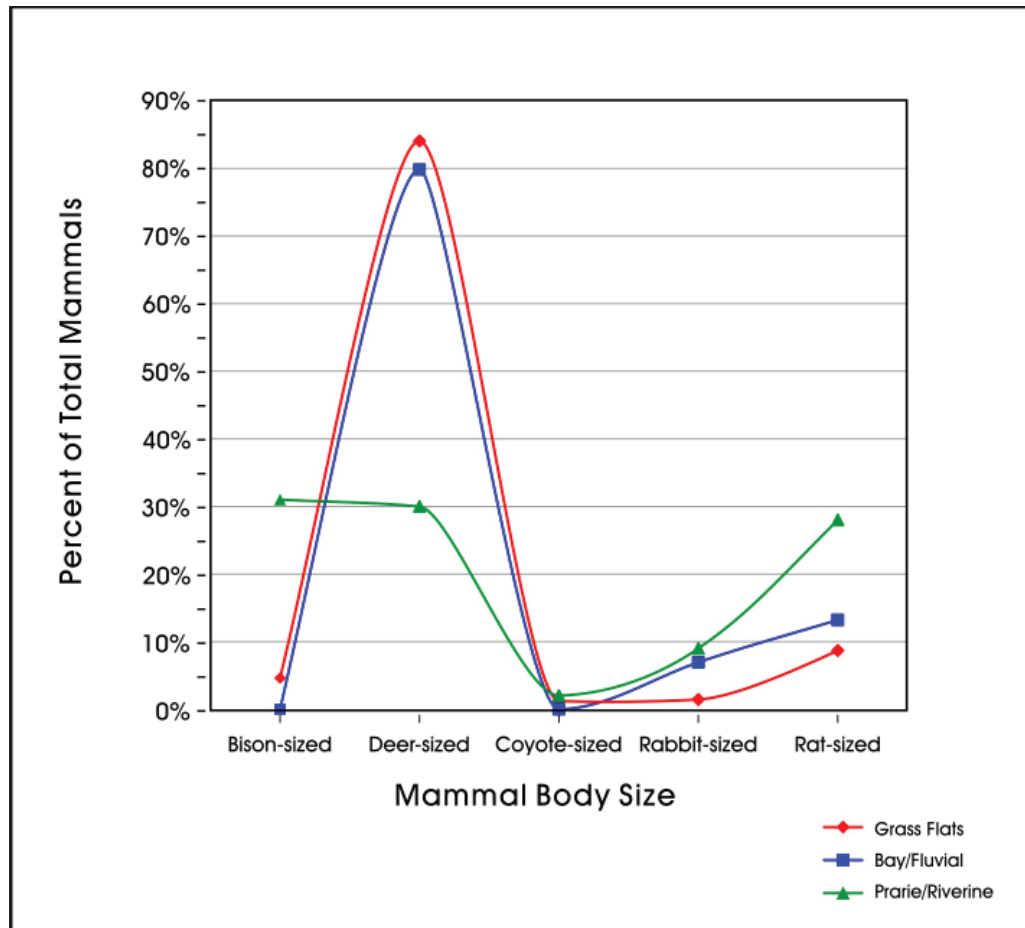


Figure 11-2. Comparison of the relative utilization of mammals in five body-size categories in sites in the three habitat types.

were only occasionally available near the bays, and rare or never found near the Grass Flats sites. In general, however the people inhabiting the Bay-Fluvial habitat concentrated on deer, to the near exclusion of smaller mammals, or, indeed, reptiles or birds (see Figure 11-1). Given that the amount of meat represented by each bone is larger in Large mammal bones than in the smaller categories, this emphasis on deer can be presumed to be even large than Figure 11-1 indicates, though the relative importance of Very Large mammals should be increased as well. In Bay-Fluvial sites, it appears that rabbit and rat-sized animals are replaced with molluscs in the hierarchy of resource status.

In the Reptile Category, turtles and snakes are the most prevalent, as Table 11-8 shows. By far the largest number of turtle and snake bones were found in the Grass Flat sites. Both types of reptiles would be common in the marshes, but it is a little surprising to find so many snakes. The only time when snakes would be relatively easy to catch would be on cold days, when they would have difficulty moving fast enough to get away from humans. Both snakes and turtles would be

less readily available near Bay-Fluvial sites and the numbers in Table 11-8 show this. Sites in the Prairie-Riverine habitat have many turtle bones, but few snake bones. This probably reflects the relative densities of the two types of reptiles in the riverine environment. While terrestrial turtles and tortoises, as well as snakes, would have been plentiful on the prairies, they would have been much harder to find in the large prairie setting. Finding turtles in the confines of the riverine setting would have been easier.

Fish are the final category of vertebrates. The complete list of identified taxa is available in Appendix H. For this discussion we will concentrate on the more common fish identified to at least Genus. Table 11-9 shows the eight genera of fish that had more than 25 bones in the sites of at least one of the three habitats. Surprisingly, they include Gars (*Lepisosteus* sp.) at one Grass Flats site (see Appendix H) Gars are a freshwater fish in four species, two of which are common in the rivers and larger creeks of Texas, and one of which, the Spotted Gar (*L. oculatus*) is common in the lower rivers as far west as the Nueces, and has been found in brackish water

Table 11-9. Count of Bones from Selected Species of Fish from Sites by Habitat

Site Habitat	Fish Species							Total Identified Fish
	Gars (freshwater)	Sea Catfish	Sheepshead	Sea Trout	Atlantic Croaker	Black Drum	Red Drum	
Grass Flats	28	92	18	45	3	39	14	239
Bay-Fluvial	0	126	2	349	572	529	344	1,922
Prairie-Riverine	0	22	5	13	7	20	7	74
Totals	28	240	25	407	582	588	365	2,207

(NatureServe Explorer 2008). This would be especially likely in the aftermath of large rain events, when the salinity of the bays has been lowered.

The remaining fish in Table 11-9 are all common in estuarine waters of the Gulf of Mexico today (Stickney 1984:204). These seven fish genera were consistently the most commonly identified fish in all the comparative sites. In fact the total number of fish in the comparative study that were within these seven genera was 2,249, while the total number of fish identified that were not in these genera was 68 (Appendix H). One reason for this is probably their size. By examining the data in Table 8-3, we can see that the average size of fish in the seven genera in Table 11-9 is 424.1 mm, while the average size of the 7 remaining species shown in Table 8-3 is 229.2 mm. However, it is important to note that we cannot assume that these seven genera were actually the most commonly chosen as prey by the inhabitants of the sites in this study. The large size of adults of these species probably put them higher on a food resource value list, but their large size also makes them more identifiable and makes them seem a larger portion of the fish diet than they may have been. It is possible that, as at 41NU2, large numbers of small fish were exploited, gathered in the shallow waters near the Grass Flats and Bay-Fluvial sites but remain unidentified during faunal analysis.

We are again hampered in our analysis by the “analytical absence” (Lyman and O’Brien 1987) of small fish, due to the extreme friability of their bone. Field observations at 41NU2, as noted above, showed that in some locations there were layers of fish bone representing hundreds or possibly thousands of small fish. The fragility of their bone prevented enumeration, however. Bone stable enough to survive the disturbance of collection and floating could be found with the centrums of vertebrae from about 5 mm to almost microscopic. This is no surprise. All of the fish listed in Table 8-3 spend all or the majority of their juvenile period in the shallow waters of the Corpus Christi Bay system, joined by at least nine species of fish, Order Cyprinodontiformes, that are all small fish that live year-round and are abundant in the shallow estuarine waters as well (Stickney 1984:210-211). At

some times of the year the biomass of small or even tiny fish available by scooping up in baskets must have been worth the effort, as the layers of small fish bone at 41NU2 suggest.

We are left, then, to note that the identifiable fish were all large. Like all marine fish that live part of their lives in estuaries, they can all tolerate low salinity, in most cases even freshwater. All are tolerant of at least mildly hypersaline conditions and three can tolerate extreme hypersalinity (Table 8-3). All can tolerate low water temperatures, and all can tolerate water temperatures of 35° C (95° F) or above. Their ability to withstand these extremes of salinity and temperature, at least for a while, is necessary, given the extreme variability of these factors in estuaries, especially estuaries like the Corpus Christi Bay system, which is fed only by rivers and creeks running through arid regions which are not infrequently deluged by massive rain events.

The two most common species are the Black Drum (*Pogonias cromis*) and the Atlantic Croaker (*Micropogonias undulates*), possibly because these two fishes are especially tolerant of both low and high salinities, and low and high water temperatures (Table 8-3) making them potentially available even during weather extremes.

Invertebrates

A total of 38 different mollusc genera were identified in the 17 sites in the comparative database (see Appendix H). However, as mentioned at the beginning of this study, the expertise of the faunal analyst in the reports that provided the data used here was variable, as was the interest in smaller shells that were unlikely to have been deliberately gathered for food. Again, all we can do is make use of the data we have, keeping in mind this caveat.

The species include very small molluscs that were probably not directly collected for food, as well as large clams or gastropods. The land snail of the *Rabdotus* genus is not included in this study, since, while most reports mention the

presence of this snail, only a few enumerate them, making it difficult to know if the large numbers recovered from 41NU2 was an anomaly or a common phenomenon.

Species richness in the three habitats within which the sites are located is variable, with the Bay-Fluvial sites having 37 genera, while the Grass Flats have 18 genera and the Prairie-Riverine sites have only 10 genera. This variability is very likely the result of availability. The narrow range of habitat conditions favorable for mollusks in the Prairie-Riverine habitat is expected to suppress species richness. In contrast, the variety of estuarine habitats available to most of the Bay-Fluvial sites would include a wide variety of potential prey.

Table 11-10 is a breakdown of 19 mollusc genera recovered from the sites in the comparative database. Species smaller than the Crossbar Venus clam are excluded, as they are all very small and we assume not to have been systematically harvested for food. It is interesting that not one of the genera presented in Table 11-10 can be found in even roughly similar numbers across all three habitats. In fact six genera appear in only two habitats and five were found in only one habitat.

In order to have a better understanding of what estuarine habitats were being exploited by the inhabitants of the sites in

the comparative database, we grouped the sites within habitat categories and assigned the mollusks in Table 11-10 to an estuarine habitat. Figure 11-3 shows where the inhabitants of the sites in each habitat concentrated their collection of mollusks. The mollusks from Prairie-Riverine habitats are, not surprisingly, almost exclusively from a brackish to riverine environment. The inhabitants of the Grass Flats sites concentrated on genera found around oyster reefs and, to a much lesser degree, mudflats. The Bay-Fluvial site inhabitants made extensive use of Oyster Reefs and Seagrass Beds, with lesser use of mudflats.

Note that the three genera listed as inhabiting Oyster Reefs, Mudflats, and Seagrass Beds are all predatory gastropods that can be found hunting their prey in these three environments. The presence of so many Lightning Whelks in most Bay-Fluvial sites is probably a reflection of the shell fishing concentrated in the three environments that are also the hunting grounds for these common carnivorous snails.

This exploitation pattern suggests that hunter-gatherers tended to focus on estuarine habitats that were closest to their sites for harvesting mollusks. The data seem to indicate that the transportation of mollusks across habitats was not practiced or if it was, it occurred after some field

Table 11-10. Counts of Marine Shells from 17 Sites Grouped by Habitat

Taxa	Estuarine habitats	Site Habitat			Total
		Bay-Fluvial	Grass Flats	Prairie-Riverine	
<i>Amygdalum papyrium</i>	Seagrass flats	90	0	0	90
<i>Argopecten irradians</i>	Seagrass flats	8,287	387	0	8,674
<i>Atrina serrata</i>	Mudflats	99	588	0	687
<i>Brachidontes</i> sp.	Oyster Reefs	3	0	0	3
<i>Chione cancellata</i>	Oyster Reefs	1,541	130	1	1,672
<i>Crassostrea virginica</i>	Oyster Reefs	17,427	10,545	491	28,463
<i>Dinocardium robustum</i>	Mudflats	80	0	3	83
<i>Ischadium recurvum</i>	Oyster Reefs	100	47	71	218
<i>Littorina irrorata</i>	Saltmarshes	101	5	1	107
<i>Macrocalista nimbosa</i>	Mudflats	3,075	111	2	3,188
<i>Mercenaria campechiensis</i>	Oyster Reefs	935	16	0	951
<i>Noetia ponderosa</i>	Oyster Reefs	75	0	0	75
<i>Ostrea egestris</i>	Oyster Reefs	0	160	0	160
<i>Rangia</i> sp.	Brackish to fresh	277	9	14,608	14,894
<i>Busycon sinistrum</i>	Oyster Reefs; Seagrass	5,562	82	0	5,644
<i>Busycotypus</i> sp.	Seagrass	77	4	0	81
<i>Fasciolaria</i> sp.	Seagrass	260	25	0	285
<i>Pleuroploca gigantea</i>	Oyster Reefs; Seagrass	87	8	0	95
<i>Polinices duplicatus</i>	Oyster Reefs; Seagrass	621	0	0	621
Total "Food-size" molluscs		38,697	12,117	15,177	65,991

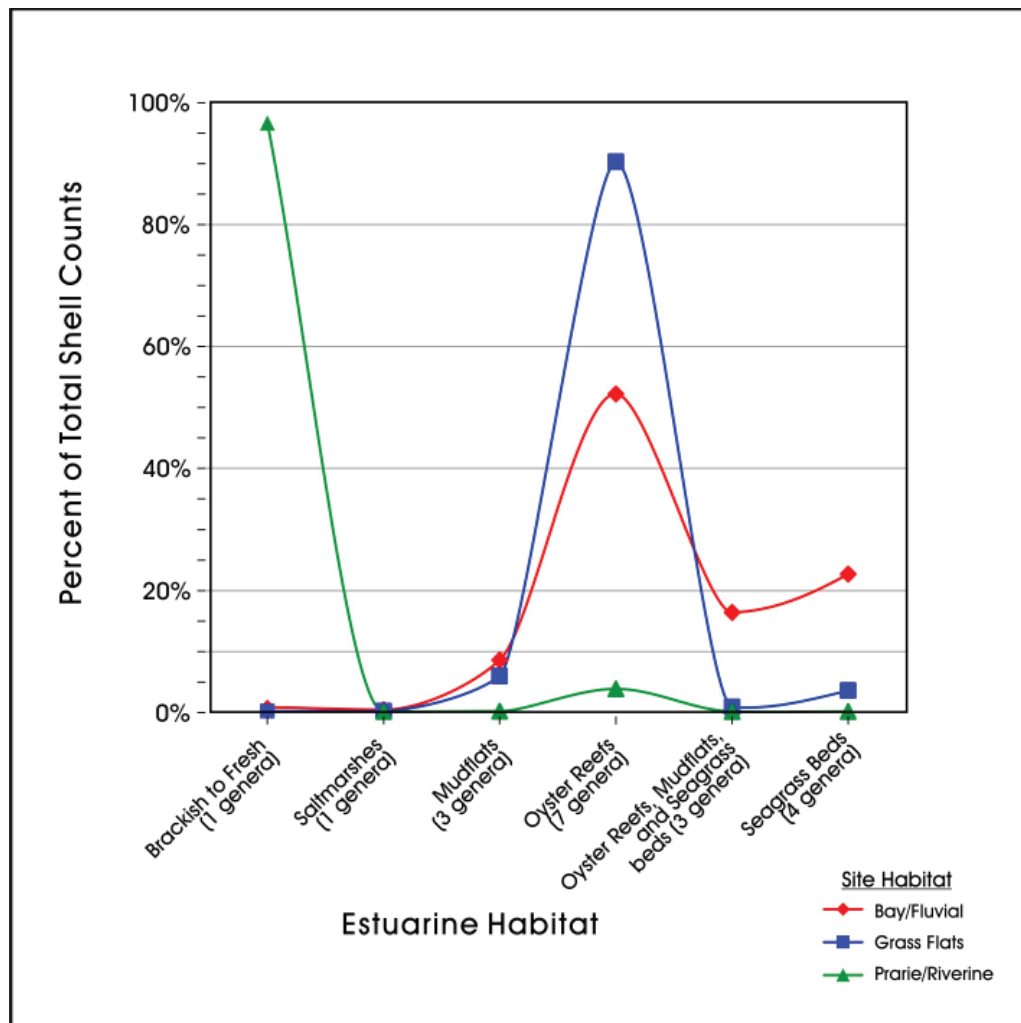


Figure 11-3. Percent of total NISP of shells in each of the three site habitat categories that were derived from each of six estuarine habitats.

processing for transportation. The later possibility is rather unlikely given the small size of most species involved and the theoretical expectation that such field-processing would most likely occur in the case of larger, heavier, species.

Comparing Use of Terrestrial, Riverine-to-Brackish, and Marine Resources

Using NISP to compare Terrestrial to Estuarine resources is highly problematic. Only using the best estimates of average calories provided per unit of weight of the remains being studied could come close to providing a reasonable estimation of relative importance in the diet. While data on the calories (and various nutrients) per gram of useable food is available or can be reasonably estimated for most species (see Chapter 8), there is no way to derive reliable estimates of the weight of useable food available to prehistoric peoples that is represented by the only materials generally available to archaeologists:

highly fragmented bones and shells. This is the perennial problem for faunal analysts (Lyman 1987, 1994; Reitz and Wing 1999): how do you estimate in a reasonable way, the dietary importance of the species you find in archaeological sites? It becomes really difficult when the faunal remains span not only Class and Order, but even Phylum.

This is not to say that there is no point in looking at the difference in NISP in the three habitats by large faunal groups. It is necessary to realize that future methods devised to increase the reliability of our interpretation of these data may someday be available, allowing much more sophisticated interpretations of available data. In the mean time we present the data we have.

Our primary interest here is to understand the relative importance of fauna collected from various ecosystems available to prehistoric inhabitants in the vicinity of

Corpus Christi Bay. Simply dividing the various taxa into vertebrate/invertebrate or mammal/fish/invertebrate will not accomplish this. It is necessary to be more specific, and assign taxa to a habitat: Terrestrial, Marine, and a third category, Riverine-to-Brackish (see Table 11-11). The latter intended to encompass those species that are most common in rivers and the boundary conditions in deltas and where the bay water is heavily influenced by freshwater input from a major creek or river. In some cases assignment of a taxa to one of these categories is not straight forward. Whether birds can be considered “Terrestrial” or “Riverine-to-Brackish” would be problematic in any case, and when so few of the bird bones were identifiable to Genus (n= 13 of 190 bird bones in entire database) it becomes even more difficult. Somewhat arbitrarily, then, all birds were assigned to Riverine-to-Brackish, based the increased likelihood of catching them when they are in riverine or deltaic settings. Turtles, though some may have been tortoises from the nearby Prairie habitat, have been placed in the Riverine-to-Brackish category as well, based on the assumption that riverine turtles are far more common

Table 11-11. Fauna Assigned to Faunal Habitat

Faunal Habitat	Fauna Assigned to Habitat
Terrestrial	All mammals, snakes, lizards
Riverine-to-Brackish	All birds, alligators, turtles, Gars, and brackish-water clams (<i>Rangia</i> sp.)
Marine	All fish except Gars, all marine molluscs

and more likely to be caught than completely terrestrial species. Table 11-11 shows what types of fauna were placed in which of these three categories.

Figure 11-4 shows the relative reliance, based entirely on NISP, of inhabitants in each of the site habitats to faunal from the three faunal habitats. From this figure we can see that inhabitants of the Bay-Fluvial and Grass Flats used the marine mollusc far more than any other fauna type, while the Prairie-Riverine depended on fauna from the Riverine-to-Brackish setting.

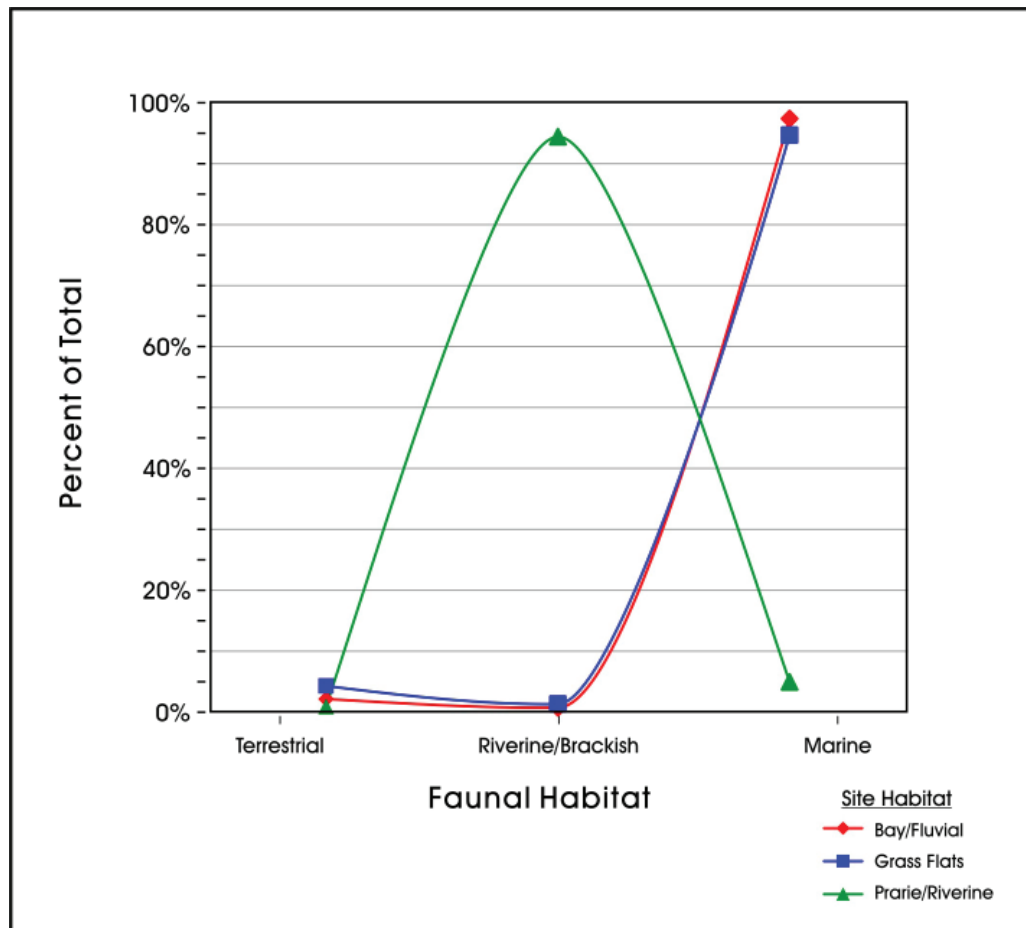


Figure 11-4. Percent of Total NISP of species found in the sites in the three habitat categories categorized into Terrestrial, Riverine/Brackish, and Marine environments

It is interesting to consider the Riverine-to-Brackish category in more detail. In Table 11-12 we can see that the inhabitants of the Prairie-Riverine sites made use of a great deal of invertebrates, just as those living at the Bay-Fluvial and Grass Flats sites did, however it was the Rangia (Rangia sp.) clam that they concentrated on. Though adults of both clams can withstand sudden changes in salinity and temperature, the Common Rangia (Rangia cuneata) needs a salinity range of 0 to 10 ‰ for the survival of larva (Andrews 1977:200-221). The Brown Rangia (Rangia flexuosa) prefers more salinity, but it too is more common in deltaic and other heavily river-influenced areas (Andrews 1977:200-221). Thus the inhabitants of the Prairie-Riverine sites were not as different in the kinds of foods they ate as Figure 11-4 suggests.

They, too, ate large quantities of molluscs with other faunal resources making up a small percentage of the NISP.

Table 11-12. Distribution of Fauna from Riverine-to-Brackish Habitat by Site Habitat

Fauna from Riverine-to-Brackish Habitat	Site Habitat		
	Bay-Fluvial	Grass Flats	Prairie-Riverine
Birds	110	69	11
Alligators	1	1	2
Turtles	1	140	137
Fish (gars)	0	28	0
Invertebrates (Rangia sp.)	277	9	14,608
Totals	389	247	14,758

Chapter 12: Regional Land Use Patterns

Jennifer A. Thompson, Leonard Kemp and Barbara M. Meissner

The Regional Land Use and Subsistence Model

In Chapter 9, we defined a series of expectations regarding hunter-gatherer regional land-use patterns through space and time in the five ranked habitats in the Coast Plains and Marshes region.

Briefly, we proposed that when bison are present only in low numbers or are absent (i.e., latter part of the Early Archaic and Middle Archaic periods) on the Coastal Plains and Marshes the Prairie-Riverine habitat would be ranked highest in terms of potential intake rates based on high population densities of deer. Based on the habitat choice model, because the Prairie-Riverine habitat offers the highest potential intake rate we suggested that the earliest hunter-gatherer populations coming into the region would settle in this habitat. Therefore, we would expect that some of the oldest traces of occupation within the region (i.e., residential campsites) would be found within the Prairie-Riverine habitat.

Daily foraging trips would be contained within the river valley and may access resources upstream and downstream as well as at the ecotone between the woodlands and the grassland prairie. During the early stages of Prairie/Riverine habitat occupation, subsistence practices would be focused on the hunting of deer from residential camps during daily foraging trips. Given high return rates and reduced transportation costs, no field processing of the carcass would be expected. As intra-habitat return rates decline and foraging trips increase in length, field processing may be adopted with low meat-utility elements being discarded in the field rather than returned to residential camps. Low ranked animal species may be included in the diet when captured by women during their daily foraging trips.

Based on the habitat selection model, we predicted that as hunter-gatherer population densities increased, friction between groups would lead to resource depletion and the widening of the diet breadth. As lower ranked resources are incorporated into the diet, neighboring habitats may provide higher encounter rates for lower ranked prey types if they represent the prey's preferred habitat. Initially, adjustments in resource procurement may simply consist of increases in daily foraging trip lengths and greater residential mobility (e.g., both the length and frequency of moves). Logistical mobility may also be employed to reach more distant but more profitable resources such as deer or pronghorn. Eventually, to more efficiently exploit the broader range of lower ranked

resources, human populations may begin to shift residential foraging to the new habitat.

Due to the low suitability of the prairie habitat, we proposed that although it is the second highest ranked habitat, it would not be the focus of settlement. Rather, we suggested that human populations would reorient their subsistence focus to the next lower ranked habitat, namely the beach-fluvial interface (i.e., the marsh). The exploitation of the beach-fluvial interface habitat would most likely have occurred from the nearby Prairie-Riverine setting because marshy areas are rarely suitable for human habitation. The diet of groups inhabiting the beach-fluvial interface would likely include a preponderance of lower ranked prey species some of which could be effectively procured en-mass with rudimentary technologies (i.e., technology that does not have high energy costs to manufacture) and would entail minimal or no processing costs. It is likely that the prairie would continue to be exploited through logistically organized mobility to target highly ranked resources.

The Bay-Estuary-Lagoon or Grass Flats habitat, the fourth ranked among the five coastal habitats, would be accessed from the nearby shores and habitation would have centered near the mouths of freshwater intermittent creeks that emptied into the lagoons and provided a seasonal source of water. The shores of ephemeral ponds and seasonal lakes may also have attracted occupation. As in the case of the Beach-Fluvial habitat, the availability of migratory birds and schooling fish could have led to increased encounter rates and to highly profitable bulk procurement episodes.

The lowest ranked habitat in the region is the beach and beach-proximate zone. Nonetheless, because it is near the Bay-Estuary-Lagoon habitat, it could have served as staging grounds for short procurement trips into the estuaries or processing stations for resources identified for transport. Field processing would rarely be expected since it would not be profitable to spend energy field-processing most lower ranked resources.

Regional Land Use Patterns

To examine regional land use we created a database containing information on a four-county region surrounding 41NU2. It was built from data gathered from the Texas Archeological Sites Atlas. The four counties include Refugio,

Aransas, Nueces, and San Patricio. The data collected on all prehistoric and protohistoric sites includes information on site size, age of components, artifact assemblage content, and distance to water when noted on the site survey forms.

The site locations were brought into a GIS database and mapped onto regional maps showing the distribution of habitats across the four county area (see Figure 2-12a-b). Once geo-referenced, we defined the boundaries of the different habitats found across the region (the Prairie, Prairie-Riverine, Beach-Fluvial Interface or Estuary, and Beach or Grass Flats) (Figure 2-13) and calculated their aerial extent. The boundaries of the Prairie-Riverine habitat were artificially set at 500 meters on either side of the course of the stream.

When temporal assignments were present either on the Survey Forms or the technical reports we incorporated them into the database without changes. When no temporal affiliation was provided for a site or component, we reviewed the projectile points recovered and assigned a time period, if feasible.

The four-county GIS dataset includes 553 sites: 97 from Aransas County, 266 from Nueces County, 18 from Refugio County, and 172 from San Patricio County. Figure 12-1 shows the distribution of the sites across the four-county region. It is evident that by far the largest habitat is the Prairie. The Beach-Fluvial interface (i.e., the marsh) is a distant second followed by the open bay. The Prairie-Riverine habitat is not substantially larger than the beach and grass flats habitats.

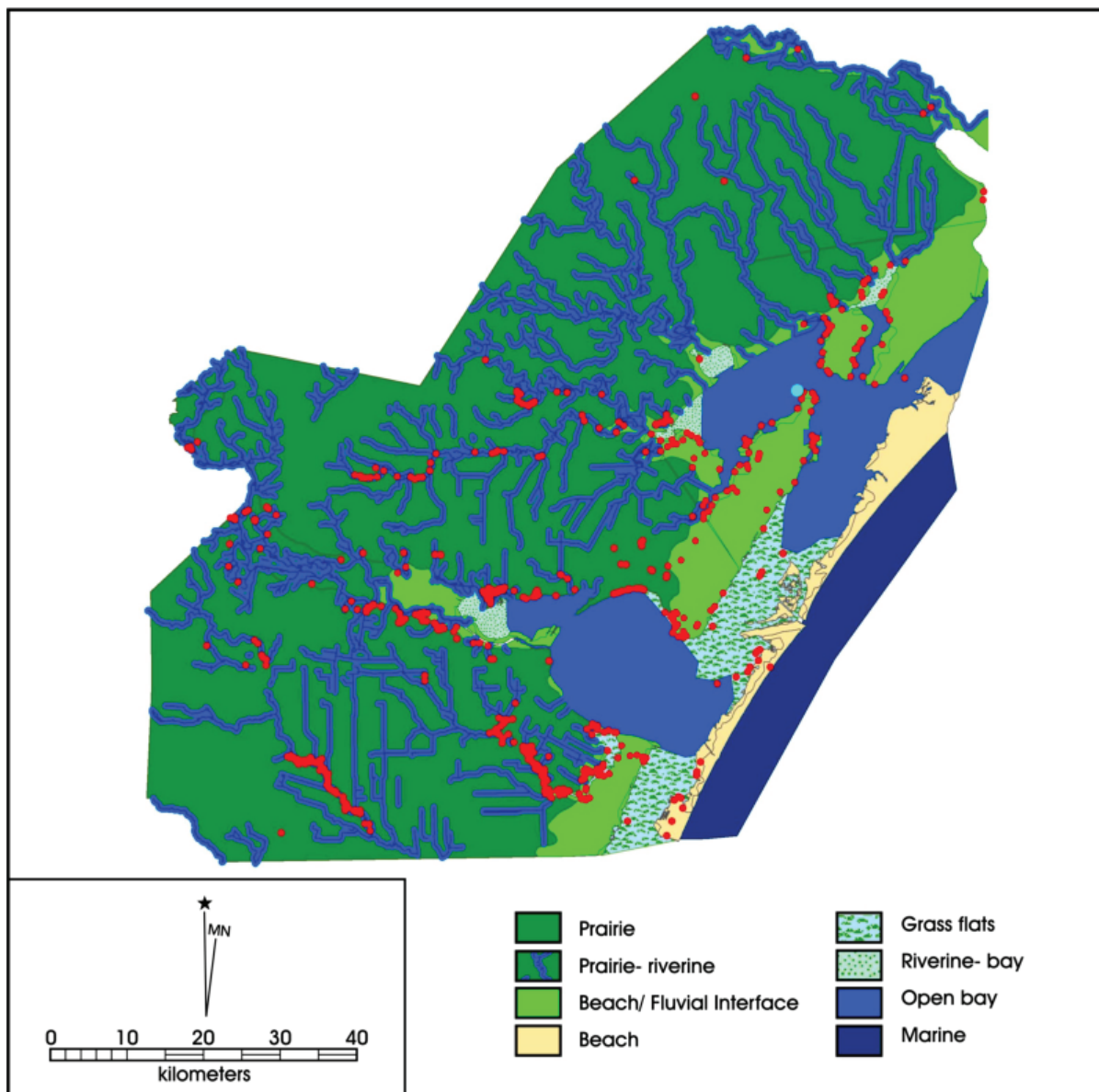


Figure 12-1. Site distribution across habitats in Aransas, Nueces, Refugio and San Patricio Counties.

Although it is the third smallest, the Prairie-Riverine habitat contains the highest site density with nearly one site per square kilometer. Among the three terrestrial habitats, the Prairie has the lowest site density. For obvious reasons, site densities are very low in the grass flats and open bay although the presence of some sites that had been recorded while water levels were very depressed does indicate that these areas were exploited when conditions allowed.

The low density of sites on the prairie may lend support to the model's assertion that the habitat was not well suited to habitation. In contrast, the fact that the highest density of sites is found in the Prairie-Riverine habitat seems to agree with the model's expectation that given its high potential return rate it will be the focus of human occupation of the region. Interestingly, the Beach-Fluvial interface (i.e., the margins of the marshes), contain the second highest site density. This finding seems to again corroborate the model's expectation that outside of the Prairie-Riverine habitat beach-fluvial settings may be the most actively utilized as stations from which to exploit the resources of the Bays-Estuarines and Lagoons.

While these patterns may accurately reflect prehistoric land use practices, they may also be a product of sampling bias. That is, river valleys tend to be some of the most intensively surveyed environments because they are considered to be high-probability settings for archeological sites. The calculation of site densities by survey areas size is one way in which such a bias may be gauged. Unfortunately, such data was not readily available for the four-county region and therefore these calculations could not be performed.

The four-county database also allowed us to explore in greater detail spatial relationships of sites to freshwater sources. As an aspect of habitat unsuitability and to highlight the challenges faced by hunter-gatherers in utilizing habitats outside of the Prairie-Riverine zone, we have already highlighted the critical importance of freshwater resources. The four-county GIS database allows us to actually quantify this aspect. Table 12-1 shows that by definition the shortest distance from sites to freshwater is found in the Prairie-Riverine habitat. While the prairie lacks localized freshwater resources, the sites found within this habitat are on the average no more than 1915 meters from the nearest stream. This two kilometer distance

is less than the typical daily foraging round trip distance of 6-8 kilometers. The shorter distance may be in response to the added costs of transporting water to sustain foraging or short term logistical forays onto the prairie margins. Table 12-1 also shows that sites on the Beach-Fluvial habitat are on the average nearly five kilometers from freshwater and the beaches fronting the bays and lagoons place hunter-gatherers nearly 14 kilometers from the nearest source of freshwater.

It is worth noting here that freshwater is defined as a flowing stream and we recognize that many localized sources of fresh or brackish water would likely have been available to hunter-gatherers exploiting resources in the Beach-Fluvial habitat or on the beach. These sources would have consisted of ox bow lakes, ponds such as the one north of 41NU2, playa lakes, and small ephemeral drainages that carried runoff during the rainy season. The Beaumont Clay underlying parts of the landscape itself could have contributed to the pooling of water across the landscape. So, access to freshwater may not have been as major a challenge as suggested by the figures in Table 12-1. It is nonetheless true, that while localized freshwater sources may perhaps been available during the wet season, during seasonal periods of low rainfall, between late June and mid-August and again between November and March (see Figure 2-3), coastal populations would have had a more difficult time securing freshwater in these coastal habitats particularly when the ponds dried out. The diatom research on Core 2 from the pond does indicate that particularly during the later part of the paleoclimatic sequence the pond may have experienced seasonal drying (see Winsborough, Appendix D this volume). Such seasonal diminishing of freshwater across these coastal habitats would have severely limited the exploitation of the beach-fluvial and estuary resources to no more than a few days at a time.

To further explore trends in the regional site distribution data, we examined patterns in site sizes across the habitats. Of the 553 sites in the database, 337 (60%) have data on site size. Table 12-2 presents the breakdown of sites by size groupings within the regional habitats. Due to small sample sizes we combined the beach and grass flats habitats into one group. Also, to insure reasonable sample sizes in most of the cells of the table, we combined site sizes into four rather large groupings. We also calculated the adjusted residuals for the table (see Everitt 1977; Haberman 1973). Adjusted residuals

Table 12-1. Geographical Habitat Information

Attribute	Prairie	Prairie-Riverine	Beach-Fluvial	Beach	Grass Flats	Open Bay
Area (km ²)	5462.3	288.32	952.37	215.64	279.1	790.67
Number of sites	135	239	145	9	25	15
Site density	0.02	0.83	0.15	0.04	0.09	0.02
Distance to Fresh Water (m)	1915	206	4790	13956	11998	7091

Table 12-2. Breakdown of Sites by Size Group within Four Habitats

Site Size Group	Prairie	Adjusted Residual	Prairie-Riverine	Adjusted Residual	Beach-Fluvial	Adjusted Residual	Beach and Grass Flats	Adjusted Residual	Totals
< 1,000 sq m	34	1.02	49	-2.51	25	2.49	6	-0.37	114
1,001-5,000 sq m	26	-0.18	58	1.18	13	-0.76	4	-1	101
5,001-10,000 sq m	9	-1.22	31	1.99	5	-0.93	2	-0.53	47
> 10,000 sq m	20	0.06	39	-0.1	8	-1.22	8	1.97	75
Totals	89		177		51		20		337

* residuals that are higher or lower than +/-1.96 are statistically significant at the .05 level

determine the contribution of each individual cell within a contingency table to the overall significance of the table. That is, when a chi-square value is significant, it is often due to the patterns within a limited number of the cells of a table. A deviations of +/-1.96 from expected frequencies signals a statistically significant divergence from expectations at a probability beyond the .05 level. The table indicates that the counts of sites diverge from expectations based on sample sizes in four cells. Sites smaller than 1,000 m² are significantly under-represented in the Prairie-Riverine habitat while sites measuring between 5,001 and 10,000 m² are over-represented. That is, fewer small sites are found than expected and a greater number of large sites are present than expected based on regional site size patterns. In the Beach-Fluvial habitat, small sites are over-represented in contrast to the Prairie-Riverine zone. Finally, in the Beach and Grass Flat habitat, very large sites measuring more than 10,000 m² are over-represented although sample sizes are rather small overall.

The under-representation of small sites within the Prairie-Riverine habitat may be due to the relative abundance of resource. That is, relatively higher resource abundance allows longer-term occupations and also potentially more frequent reoccupations of sites due to more rapid regeneration of plant communities. The repeated reoccupation of terraces would result in gradually increasing site sizes and an overall increase in the frequency of large sites, exactly the pattern noted in the table. Occupation of the margins of the Bay-Estuary-Lagoon habitat to facilitate the exploitation of the abundant but lower ranked resources may explain the over-representation of small sites within the Beach-Fluvial habitat. In addition, the pattern may also be due to the episodic and short-term nature of the visits to this habitat to exploit resources in an opportunistic manner. Given that the lack of freshwater resources does not provide a settlement anchor that would favor the re-occupation of specific spots in the habitat, the use of the Beach-Fluvial zone may be more extensive depending on which part of the Bay-Estuary-Lagoon provides an abundance of resources at the time of the visit. Finally, the over-representation of large sites in the combined Beach and Grass Flats habitat may reflect the fact

that dunes on the beach may represent settlement magnets that would be reoccupied on a systematic basis leading to increases in site size over time. Alternatively, large sites may not only be confined to the dunes but may also result from the intermittent use of different parts of the beach over time. These occupations can eventually coalesce into what appear to be low density material distributions spread over large areas. The key difference between these large sites and those in the Prairie-Riverine habitat is that the coastal ones are a product of the slow merging of small intermittent reoccupations spread linearly along the beach. In the case of the Prairie-Riverine habitat, the abundance of resources actually encourages re-settlement of productive spots on the landscape.

In the research design developed for this project, we suggested based on the Habitat Choice Model that the habitat that would show the earliest signs of occupation would be the Prairie-Riverine zone. Only after resource depletion and the gradual reduction of highly ranked resource from the habitat would hunter-gatherers begin to move out into neighboring habitats with similar rankings.

To explore temporal changes in settlement across the coastal zone, we compared the ages of components at each site in the GIS dataset across habitats to assess the expectation that the oldest sites in the region are in the Prairie-Riverine habitat. Of the 553 sites in the database, only 214 components have temporal affiliations defined in broad terms (i.e., Paleoindian, Archaic, or Late Prehistoric). Only 74 components have information that allows them to be dated to finer-scaled temporal units (i.e., sub-periods) such as Early, Middle and Late Archaic and Initial and Terminal Late Prehistoric. Table 12-3 presents the breakdown of sites within the habitats identified by temporal affiliation.

Looking at the coarse-grained temporal divisions (i.e. Paleoindian, Archaic and Late Prehistoric), the oldest sites within the dataset are in the Prairie-Riverine habitat. Given the geological history of the coast, however, it is likely that many other Paleoindian sites are found currently submerged

Table 12-3. Components per Habitat Through Time

Time Period	Primary Habitat						Grand Total
	Prairie	Prairie-Riverine	Beach-Fluvial	Beach	Grass Flats	Open Bay	
Early Archaic	2	4	1	0	1	1	9
Middle Archaic	4	6	3	0	2	1	16
Late Archaic	11	22	8	0	1	0	42
Initial Late Prehistoric	0	3	0	0	0	0	3
Terminal Late Prehistoric	0	3	0	1	0	0	4
Archaic	23	54	15	2	2	1	97
Late Prehistoric	23	53	25	4	6	3	114
Paleoindian	0	2	1	0	0	0	3

on the continental shelf. Therefore and given the small sample size, the Paleoindian settlement, while being the earliest, it is not representative of the overall coastal settlement during the period. The pattern that is perhaps more meaningful is that beginning with the Early Archaic sub-period, and continuing into the Late Archaic, the number of sites in the Prairie-Riverine habitat are consistently higher than in any other single habitat in the region. This pattern suggests that throughout the earliest intensive occupations of the habitats examined, the Prairie-Riverine habitat was the one that was the focus of settlement and exploitation. Again, the possibility that the pattern is the product of more numerous surveys within the habitat cannot be ruled out.

Even more telling is the pattern that during the Late Prehistoric Period there is a virtual explosion in the use and exploitation of resources (habitats) outside of the Prairie-Riverine habitat (see Table 12-3, see sites assigned to a broad Late Prehistoric affiliation). While the initial stages of this reorientation in settlement and resource extraction strategy may be noted during the Late Archaic (Table 12-3, see sites assigned to the Late Archaic sub-period), the shift becomes even more amplified during the subsequent Late Prehistoric. Interestingly, the exploitation of the Prairie habitat is consistent through the Archaic and Late Prehistoric Periods.

The Late Archaic and Late Prehistoric reorientation in settlement and resource extraction strategies agrees well with expectations derived from recently evaluated models of intensification of resource extraction by hunter-gatherers in Texas (Johnson and Hard 2008). The population growth rate model considered by Johnson and Hard (2008:Figure 4) indicates that the Lower and Central Coast of Texas should be locations where under packed conditions intensification of extraction is highly anticipated. Furthermore, Johnson and Hard (2008:143) expect that it is exactly on the Lower and Central Texas Coast, where the modeled value of dependence on aquatic resources exceeds 30 percent, that prehistoric hunter-gatherer groups should turn to the intensification

of aquatic resources rather than terrestrial resources. Unfortunately, we do not have population estimates that can inform us about rates of population growth and packing. Nonetheless, the regional settlement pattern data reviewed above does suggest that, in concordance with the expectations of the Habitat Choice Model, the Late Archaic and Late Prehistoric reorientation in settlement pattern from the Prairie-Riverine to more beach-proximate habitats may be indicative of the aquatic intensification expected based on the population growth model.

What the Faunal Data Shows

In Chapter 11, we discussed our reconstructions of regional subsistence practices based on the faunal data compiled from eighteen excavated sites in San Patricio, Refugio, Nueces, and Aransas Counties. Clearly, this information also has some bearing on general patterns of land use within the broader region. Specifically, we noted that inhabitants of sites located in the Grass Flats, Bay-Fluvial and Prairie-Riverine habitats exploited a similar set of vertebrate subsistence resources (see Table 11-8) and they all utilized estuarine resources (see Figure 11-3). Nonetheless, some degree of resource emphasis is evident. For example, inhabitants of sites located in habitat use resources from fewer of the estuarine habitats (see Figure 11-3) than populations living in greater proximity to the estuaries. Furthermore, it is the brackish to fresh water habitat that is the one most often exploited from sites in the habitat. A similar pattern is evident in terms of faunal habitats exploited (see Figure 11-4). The inhabitants of sites located in the habitat tend to focus on faunal resources found in the riverine/brackish habitat and the exploitation of fauna in the marine and terrestrial habitats appears to be de-emphasized.

In summary, while a range of resources representing terrestrial, riverine and marine habitats, are being consumed by populations inhabiting the, Grass Flats and Bay-Fluvial habitats, there appears to be an emphasis on the consumption of more localized resources in greater proximity of each site

within each habitat. This general observation is supported by the study of bone stable carbon and nitrogen isotopes derived from a series of Texas prehistoric cemeteries (Johnson and Hard 2008:145-146). The data show clear divisions in isotope values between inland, coastal and riverine populations. While the distinctions in isotope values between these populations are somewhat surprising, they do suggest that these populations to a large extent focused their subsistence strategies and dietary intake on resources found within their immediate vicinity.

Conclusions

Our analysis of regional settlement patterns began with a look at site densities across the different habitats present within the four-county region encompassed by Nueces, Refugio, Aransas and San Patricio counties. While intriguing, the highest density of sites in the Prairie-Riverine habitat may be a product of the greater number of archeological surveys conducted in this habitat driven by more intensive regional developments. Unfortunately, while the number of sites by acre surveyed could provide an unbiased estimate of site

densities by habitat, we could not gather this data in time for this report. Based on the data we were able to collect, it appears that through time the Prairie-Riverine habitat maintained the highest density of sites in the region. This is an expected pattern given that it was the habitat with the highest potential intake rate for hunting and gathering. Based on the Habitat Choice Model, only when the highest ranked resources in this habitat were unavailable would we expect populations to shift to using other habitats. Identifying shifts in residential locales because of changing food availability is difficult to determine by examining site distribution because of the small sample of sites with detailed sub-period level temporal affiliation and faunal data. Nonetheless, the regional distribution of sites within the Prairie-Riverine setting, vis a vis other habitats, between the Early and Late Archaic sub-periods does appear to support the expectation that the Prairie-Riverine setting was the focus of settlement and extractive activities. Similarly, the broadening of site distributions during the Late Prehistoric to include sites in all other habitats outside of the Prairie-Riverine setting seems to suggest that by this time lower ranked resources were regularly incorporated in the diet of hunter-gatherers.

Chapter 13: Territoriality and Cemeteries

Steve A. Tomka, Jennifer A. Thompson and Leonard Kemp

Volume I of this series summarized the explorations of the prehistoric cemetery at 41NU2 (Jackson et al. 2004). The stratigraphic distribution of the burials in conjunction with the temporal diagnostics recovered suggests that the cemetery was utilized during the Middle Archaic and Late Archaic sub-periods and Late Prehistoric Period. The cemetery at 41NU2 is only one of a number of sites containing human remains across South Texas and the Central Coastal Plains. To provide a broader context within which to interpret the significance of the cemetery at the site, this chapter discusses the distribution of archeological sites containing human burials across a large study area from the coast to the Edwards Plateau. The discussion and review of the regional data focuses on whether the distribution of sites with human interments may be reflective of incipient territoriality among prehistoric indigenous groups or whether it is the product of re-occurring use of selected places across the landscape.

Theoretical Background and Texas Patterns

There is extensive anthropological literature discussing various case studies of territoriality among hunter-gatherer groups (Bishop 1970; Blundell 1980; Cashdan 1983; Dyson-Hudson and Smith 1978). These researchers define territoriality primarily as limiting access to an area in order to control the exploitation of one or more resources (see Cashdan 1983:47). Several factors appear to condition when territoriality emerges and what form it takes.

Territoriality usually results in some benefit to the group, and most often this benefit will be increased availability of resources or increased foraging efficiency. However, these increased benefits also come with costs. The defense of a territory is energetically costly and the larger the territory the greater the costs of its defense. The density of resources affects the size of the territory being controlled. So that, the more abundant the resource, the smaller the area that needs to be defended, and concomitantly, the scarcer the resource becomes the greater the territory that will need to be defended to provide sufficient intake rates for survival. The predictability of the resources also affects tendencies for territoriality or ownership of resources. Usually, resources that tend to fluctuate wildly from year to year will not be defended but those that reoccur in a predictable fashion are defendable. Population density also interacts with territoriality to the degree that it has an effect on resource availability. Competition between populations will increase as resources become scarce. In general, dense and predictable resources

will be the target of territorial defense while unpredictable scarce resources will not.

Territoriality and the control of resources can range from particular families controlling access to patches of dense oak forests and even individual oak trees (M. Anderson 2005:133-134), and salmon fishing territories along streams (Ames 1994). According to Cashdan (1983:349), territorial control can take at least two forms: (1) the defense of the perimeter or boundary of the territory, and (2) the control of access through social boundary maintenance. That is, control can be achieved through physical exclusion or social exclusion. Perimeter or boundary defense is usually accompanied by the physical marking of the boundaries of a territory, or actual patrolling of the boundary or both. In either case, the defense of the territory tends to be relatively energy intensive either in terms of the energy expended in the actual defense or the construction of signals to ward off potential trespassers. Social boundary maintenance is practiced through marriage and other social means. The territorial markers for this type of control are much more subtle and typically involve natural features of the environment rather than constructed markers (Cashdan 1983:50). Bicchieri (1969) suggests that perimeter defense often develops in situations of resource depletion and geographical circumscription (Cashdan 1983:57). According to Cashdan 1983:63), perimeter defense should develop when resource abundance is high and competition is high while social boundary defense should be practiced when resource abundance is low but competition is high. Under low levels of competition, territoriality should not develop regardless of resource abundance.

In summary, discussions of territoriality identify two dimensions to the issue: (1) the types and characteristics of the resources targeted, and (2) the conditions that lead to territorial behavior. In terms of the first, resources that are highly productive and occur in reasonably well defined or localized patches are often the targets of access control. In terms of the second aspect, control of access to resources tends to emerge when populations tend to be spatially circumscribed leading to competition for limited resources potentially resulting in the depletion of critical resources.

In addition to the factors that may lead to territoriality it is also worth considering the factors that lead to the formation of cemetery sites. Be they modern, historic or prehistoric, it is relatively clear that cemeteries are the cumulative products of a series of short-, medium- and long-term

decisions and processes. In discussing hunter-gatherer burials in southeastern Australia, Littleton (2007) defines the following short term actions associated with burial: the actual interment, and the period of mourning (Littleton 2007:Table 2). These actions remain in the consciousness of the relatives for approximately a year following the interment. Knowledge such as burial orientation and position, and visible markers of the interment such as a monument may typically be preserved for 1-4 generations, while landscape formation and climatic forces that may alter the look of a particular place beyond recognition or change the resource structure in its vicinity may take longer periods, unless they occur in a catastrophic manner (i.e., floods, earthquakes). The critical point is that cemeteries are likely to be the product of long-term processes that may take thousands of years during which time the utilization of the cemetery may not necessarily be continuous. We tend to view cemeteries from a contemporary ethnographic underpinning without necessarily recognizing the methodological difficulty faced by archaeologist that attempt to relate specific death events to long term process that result in the accumulation of interments and the formation of what we as archaeologists define as cemeteries (see also Littleton and Allen 2006). Because of exactly this difficulty, several archeologists have begun to reframe the interpretive perspective of cemeteries as not necessarily representing signals of territoriality but rather “persistent places.” Schlanger (1992:92) defines persistent places as cultural features that structure the use and reuse of landscapes. Persistent places may be formed through three processes including recognition of unique qualities of locales suited for certain activities, the accumulation of remains, features and facilities that attract reuse and reoccupation, and through long-term occupation and/or revisitation (1992:92). From the perspective of site formation processes, the concept of persistent places offers an elegant explanation for the creation of cemetery archeological sites and is being used as a feasible interpretive framework in a number of contexts (Jacobs 1995, Walthall 1999).

Perspectives on Texas Cemetery Sites

With these general theoretical roadmaps we now turn to the discussion and examination of the Texas cemetery site distribution. There seems to be agreement among scholars that the development of cemeteries in this study area is an indicator of population growth and the establishment of territories (Prewitt 1981; Story 1995; Hall 1995a; Hall 1995b; A. J. Taylor 1998). Cemeteries as markers for territories could indicate a region had a high population density and low abundance of optimal resource patches. Conversely, areas lacking cemeteries may have been less populous, less productive, or were used only by relatively more mobile groups and saw less competition for resources. Two broad

regional studies of central and South Texas cemeteries have been conducted in the context of site 41LK28, the Loma Sandia site (A.J. Taylor and Highley 1995, Taylor 1998). More recently, Perttula (2001) reviewed mortuary data from thirty years of excavations of hunter-gatherer sites across the Rio Grande and Central Coastal Plains.

These reviews show similar trends and could support the idea of territories tied to the most productive resource bases and marked by cemeteries in the Late Archaic and Late Prehistoric. There is little evidence of Paleoindian and Early Archaic territoriality in the archeological record in South Texas (A.J. Taylor 1998). Cemetery sites cluster along the Nueces River and Corpus Christi Bay, the lower Pecos River, the lower Guadalupe, and the Colorado and Brazos River basins, that is, riverine setting in the Coastal Plain (Perttula 2001). These riverine setting have large cemeteries in both Late Archaic and Late Prehistoric whereas other areas only support large cemeteries in either the Late Archaic or the Late Prehistoric. The Central Texas Coastal Plains region has very few sites with burials.

A.J. Taylor (1998) identifies territoriality in part by distribution of cultural materials. After 3000 B.C., projectile point distributions begin to show some patterning, especially among the Tortugas, Nolan, and Travis types (A.J. Taylor 1998). The initial appearance of extensive burial sites in southern Texas dates to about this same time (2500 to 3000 B.C.).

Also related to territoriality studies is population density. Some interpret an increase in the amount of “habitational debris” at interior sites in South Texas during the Middle and Late Archaic (Hall et al. 1986; and Black 1989: 49, 51) to an increase in population density. Late Archaic coastal sites also show a marked increase in cultural material. A rise in site numbers and richness could be attributable in part to the establishment of the barrier islands, marshlands, and estuaries which in turn could have promoted population growth. Also during the Late Archaic, A.J. Taylor (1998) reports an increase in the number of large mortuary sites in South Texas and along the Texas Gulf Coastal Plain and views the development of extensive hunter-gatherer cemeteries as territorial markers. Hall (1998) interprets the increased reliance on aquatic resources and increase in the number of cemeteries on population pressure on the food base. He ties the distribution of territories and cemeteries to patchy but abundant resources like pecans (and other nuts), acorns, prickly pear, and aquatic resources in the coastal estuaries-bays and lagoons. Reciprocal alliances between territories allowed food surpluses in one territory to spread to others.

The Methodological Approach

Given these two dimensions of territoriality, to begin to investigate and explain the distribution of prehistoric sites with interments, we first divided the Coastal Plains and Marshes and South Texas Brush Country regions into broad habitats that were characterized by different resource availabilities. Because the scale of this analysis is larger than the four-county area we focused on in the previous land use discussion, our broad habitats are not the same as those discussed in Chapter 12. Rather, the larger region is divided into five habitats consisting of: (1) the South Texas Brush Country; (2) Prairie; (3) Oak Woodlands; (4) Oak Woodlands-Prairie Riverine; and (5) Beach-Estuary-Marsh. As much as possible, we tried to incorporate the finer-scale habitats employed in Chapter 12 in these groupings to allow for continuity (Figure 13-1). The eco-regions illustrated in Figure 13-1 are those defined by the United States Environmental Protection Agency (USEPA) under Level III and IV. There are 12 Level III eco-regions and 56 Level IV eco-regions in Texas that continue into adjacent parts of the US and Mexico. The methods for defining these regions are in the National Atlas of the United States (2003).

Based on the extensive earlier discussions of resource potential, we would anticipate that the Oak Woodlands-Prairie Riverine habitat would be the most productive of the habitats followed by the Oak Woodlands. The South Texas Brush Country and the Prairie are assumed to be less productive with the exception of linear corridors along streams that cross through them. As we described in Chapter 8, the Beach-Estuary-Marsh (same as the Bay-Estuary-Lagoon) habitat can be productive when it comes to small and very small prey that can occur there seasonally in large quantities.

Having briefly summarized our perception of the ranking of the habitats based on potential resource availability (much of it described in detail in Chapter 8), we now turn to the second dimension of territoriality mentioned above, population dynamics. Several archeologists have in the past attempted to estimate trends in population densities across Texas (Prewitt 1985; Skinner 1981; Weir 1976). These attempts focused primarily on Central Texas and for the sake of expediency we employ Prewitt's (1985) estimates as applicable to the

study region. Using the ratio of component occurrences to the duration of a component in years, Prewitt's (1985:Figure 6) estimate shows several peaks in population densities over time. The most dramatic of them occurs in the Late Archaic and is followed by a drop during the Austin Phase of the Late Prehistoric only to subsequently rise, although not nearly to Late Archaic levels.

Given the expectations derived from the general anthropological theory and based on the resource potential of the larger region and the relative population densities estimated by Prewitt (1985), we would expect that territoriality should emerge among Texas hunter-gatherers in the Oak-Woodlands-Prairie Riverine habitat. The Oak Woodlands habitat away from the river valleys may not see an extension of territorial control of its resources. A second area where territorial control may emerge is on the coastal strip, and specifically in the vicinity of the Beach-Estuary-Marsh habitat. We would expect, however, that the emergence of signs of territoriality would occur much later in time in the coastal zone compared to the Oak Woodlands-Prairie Riverine habitat. Finally, we anticipate that the hypothesized strong increase in relative population densities during the Late Archaic would spur the earliest signs of territoriality. Divergence from these expectations may suggest that the emergence of cemetery sites is not linked to the establishment of territoriality among Texas hunter-gatherers.

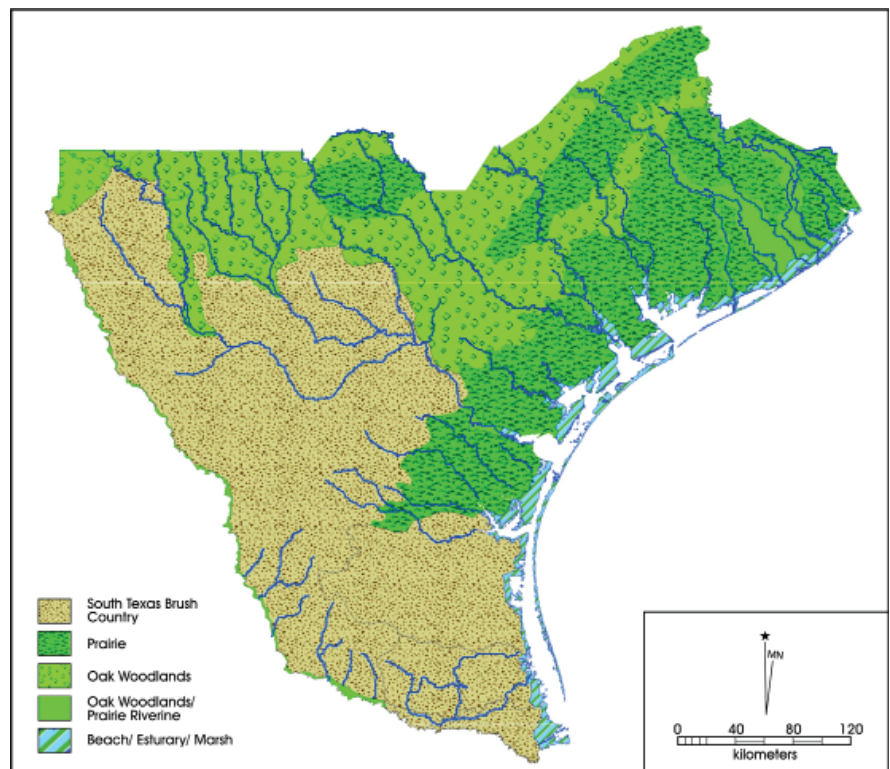


Figure 13-1. *Habitat map.*

For analytical purposes, we assume that cemeteries represent a visual signal of territoriality. It is our impression that the presence of a single interment at a site has different implications than that of multiple interments. That is, we assume that there are semiotic differences between a single human interment versus the presence of an actual cemetery (i.e., multiple human interments).

GIS Database

A mortuary database was compiled in the Geographic Information System program that includes data from the five major physiographic regions including the Gulf Coastal Plains and Marshes, Blackland Prairie, Edwards Plateau, Central Texas Plains and South Texas Plains. The database includes site name and trinomial, site size, provenience information, habitat as defined above, presence of grave goods, population estimates including the presence of multiple versus single graves, evidence of violent death, age, sex, and interment

orientation, cultural period, and radiocarbon dates. Dates and cultural period assignments are specific to the burial component; sites listed may contain components/dates not listed because no burials were associated.

Distribution of Burial Sites Across the Study Area

The study area includes portions of the Edwards Plateau, the Blackland Prairie, Central Texas Plains, the Gulf Coastal Plains and Marshes, and the South Texas Plains physiographic regions and includes 186 sites reported to have human burials (Figure 13-2, Table 13-1). The Gulf Coastal Plains and Marshes region contains almost half (n=99) of the sites with human burials. They are distributed in four habitats within this physiographic region: the Beach-Estuary-Marsh (n=25), Oak Woodlands-Prairie Riverine (n=18), Prairie (n=40), and the South Texas Brush Country (16). The physiographic

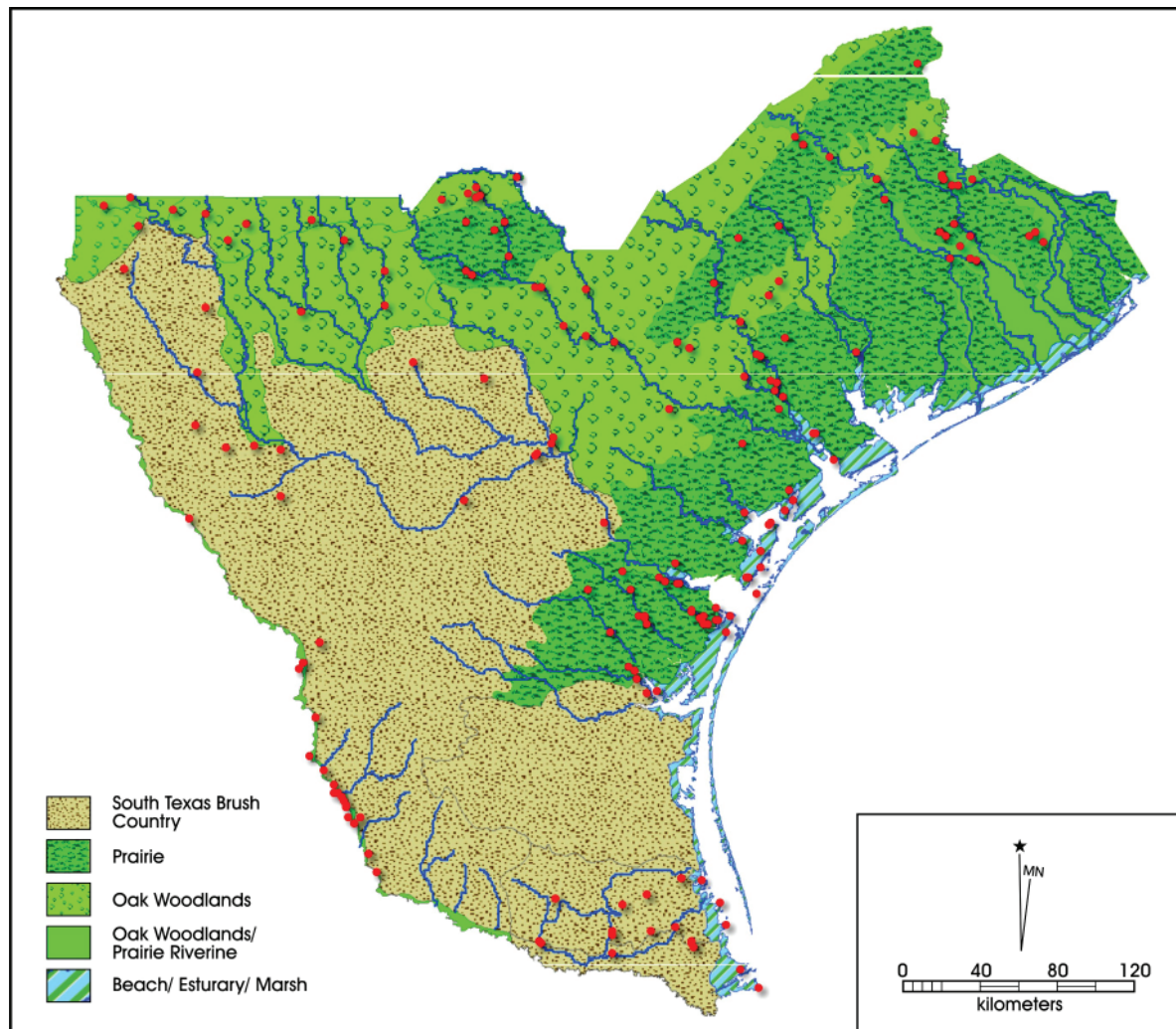


Figure 13-2. Distribution of burial sites across the study area.

Table 13-1. Counts of Burial Sites Across the Study Area

Major Physiographic Region (Gould 1975)	Habitats	Total
Blackland Prairie	Oak Woodlands-Prairie Riverine	1
	Prairie	11
Blackland Prairie Total		12
Central Texas Plains	Oak Woodlands	17
	Oak Woodlands-Prairie Riverine	2
Central Texas Plains Total		19
Edwards Plateau	Oak Woodlands	13
Edwards Plateau Total		13
Gulf Coastal Plains and Marshes	Beach-Estuary-Marsh	25
	Oak Woodlands-Prairie Riverine	18
	Prairie	40
	South Texas Brush Country	16
Gulf Coastal Plains and Marshes Total		99
South Texas Plains	Oak Woodlands	5
	Oak Woodlands-Prairie Riverine	17
	South Texas Brush Country	21
South Texas Plains Total		43
Grand Total		186

region with the next highest count of sites with human remains is the South Texas Plains (n=43), followed by the Central Texas Plains (n=19). Other physiographic regions have smaller numbers (Table 13-1).

Looking at the distribution of sites with human remains by habitats across all the physiographic regions, most human burials are within the Prairie (51) habitat (Table 13-1). Across the remaining habitats, burial sites are evenly distributed with 37 sites in the South Texas Brush Country, 38 across the Oak Woodlands-Prairie Riverine habitat, and 35 in the Oak Woodlands. The Beach-Estuary-Marsh habitat has the fewest burial sites (n=25).

These data present a rather complex picture, but at the most simplified level, it is evident from the distributions in Figure 13-2 that most sites with human interments are found along streams (i.e., the Prairie-Riverine habitat). These locations tend to be the highest ranked in terms of resource potential within each broad physiographic region. In addition, if these distribution patterns reflect territoriality, they suggest that the territories comprise portions of streams and potentially some areas along the valley margins or uplands. That is, if the distribution of sites does actually reflect group territories, the territories may consist of drainages or portions of drainages. This would suggest that the territories may have paralleled drainages and would have been oriented perpendicular to the coast.

However, before accepting these data as supporting the territorial hypothesis, it is important to recognize that the prehistoric cemeteries in this research area are characterized by relatively low visibility. That is, minimal effort that has been expended in making them serve as long-term signaling features. There were no visible markers found on surface or above ground at any of the cemetery sites prior to archeological excavations and the cemetery component was either an unexpected find or was known to exist due to the partial exposure of human remains through erosion. Sites with cemeteries occupy a variety of landforms from hills to sand dunes and stream terraces. While they tend to be more common in riverine settings (see Figure 13-2), they are by no means only found in the habitat with the highest potential return rate (i.e., the Prairie-Riverine habitat). Rather they are found in all riverine settings across a whole range of habitats suggesting that the factor that is conditioning their location is the mix of resources that are found in such settings rather than any one particular hyper-productive resource. The fact that few sites with human burials are found outside of riverine settings also supports the interpretation that localities with high rates of revisitation tend to have higher probabilities of cemetery sites compared to localities with lower rates of revisitation. Again, the higher incidences of revisitation may be conditioned by the variety of economically significant resources found in riverine settings regardless of what physiographic region a human population may be exploiting.

Distribution of Burial Sites Across the Gulf Coastal Plains through Time

Perttula's (2001) study of mortuary practices in South Texas found that burial sites in south central, inland central, and the Central Texas prairie had large cemeteries dating to both the Middle Archaic and the Late Prehistoric, while other sub-regions had large cemeteries in either one or the other of these periods. Of the 186 sites in the database, 108 sites have an assigned temporal affiliation (Figure 13-3, Table 13-2). A total of 128 prehistoric components are defined in the 108 sites in the database.

The Prairie region contains the highest number of both Late Archaic (16) and Late Prehistoric burials (16), followed by Late Prehistoric burial components in the Beach-Estuary-Marsh region (n=9), the Oak Woodlands-Prairie Riverine setting (n=7), and six sites each from the Oak Woodlands and the South Texas Brush Country. The Oak Woodlands and the South Texas Brush Country both have nine Late Archaic burials followed by four Late Archaic burials in the Beach-Estuary-Marsh region.

The fact that the highest number of dated burial sites occurs during the Late Archaic (n=48) may be significant in light of the expectation that territorial marking would be likely to occur under conditions of increasing population density. The number of burial sites dated to the Middle Archaic is very low (n=12) but it may also signal some correspondence with the peak in Middle Archaic population noted by Prewitt (1985).

Distribution of Single and Multiple Interments at Burial Sites across the Gulf Coastal Plains through Time

We also examined the distribution of single versus multiple interment sites across the region through time (Table 13-3; Figure 13-4). The number of graves, single versus multiple, per component shows even distributions among those sites with reported temporal affiliations. There are 30 Late Archaic cemetery sites with multiple burials compared to 27 for the Late Prehistoric. There are 14 and 16 sites with single graves dated to the same time periods, respectively. In contrast, ten of eleven Middle Archaic sites had multiple graves and all three Early Archaic components contained more than one grave. The proportion of multiple and single

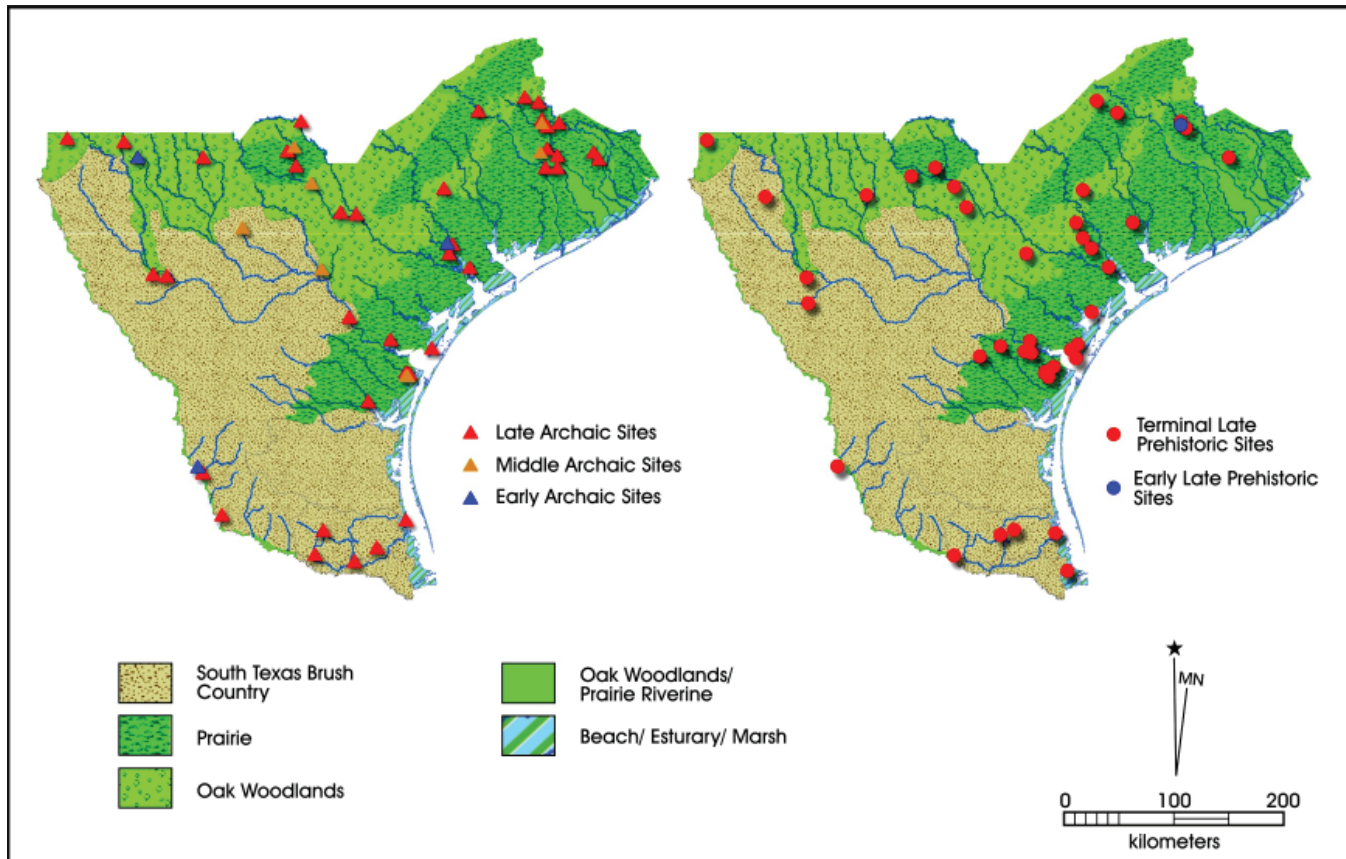


Figure 13-3. Distribution of burial sites through time.

Table 13-2. Count of Sites through Time per Habitat

Habitat	Time Period	Count of Sites
Beach-Estuary-Marsh	Late Prehistoric	9
	Early Prehistoric	0
	Late Archaic	4
	Middle Archaic	0
	Early Archaic	0
Oak Woodlands	Late Prehistoric	6
	Early Prehistoric	0
	Late Archaic	9
	Middle Archaic	4
	Early Archaic	1
Oak Woodlands-Prairie Riverine	Late Prehistoric	7
	Early Prehistoric	0
	Late Archaic	10
	Middle Archaic	0
	Early Archaic	2
Prairie	Late Prehistoric	16
	Early Prehistoric	1
	Late Archaic	16
	Middle Archaic	7
	Early Archaic	0
South Texas Brush Country	Late Prehistoric	6
	Early Prehistoric	0
	Late Archaic	9
	Middle Archaic	1
	Early Archaic	0
Total Count of Late Prehistoric		44
Total Count of Early Prehistoric		1
Total Count of Late Archaic		48
Total Count of Middle Archaic		12
Total Count of Early Archaic		3

grave cemeteries through time remains consistent during the Late Prehistoric and Late Archaic but sites with multiple interments greatly outnumber those with single burials during the Middle Archaic.

Across habitats then we would expect in general to see more multiple-grave sites than single-grave sites in each component. This remains true for all habitats except the South Texas Brush Country, where out of 16 reported sites, 12 are single-grave sites. Single-grave sites also are confined in the reported data to the Late Prehistoric and Late Archaic cultural periods in all habitats except the South Texas Brush Country where one single-grave site is reported for the Middle Archaic. Middle and Early Archaic components with multiple graves are present in the Oak

Table 13-3. Count of Interment Number through Time per Habitat

Habitat	Time Period	Multiple Interment	Single Interment	Total
Beach-Estuary-Marsh	Late Prehistoric	8	1	9
	Early Prehistoric	0	0	0
	Late Archaic	3	1	4
	Middle Archaic	0	0	0
	Early Archaic	0	0	0
Oak Woodlands	Late Prehistoric	3	3	6
	Early Prehistoric	0	0	0
	Late Archaic	6	3	9
	Middle Archaic	4	0	4
	Early Archaic	1	0	1
Oak Woodlands-Prairie Riverine	Late Prehistoric	5	2	7
	Early Prehistoric	0	0	0
	Late Archaic	7	2	9
	Middle Archaic	0	0	0
	Early Archaic	2	0	2
Prairie	Late Prehistoric	10	5	15
	Early Prehistoric	1	0	1
	Late Archaic	11	2	13
	Middle Archaic	6	0	6
	Early Archaic	0	0	0
South Texas Brush Country	Late Prehistoric	1	5	6
	Early Prehistoric	0	0	0
	Late Archaic	3	6	9
	Middle Archaic	0	1	1
	Early Archaic	0	0	0
Total Count of Late Prehistoric		27	16	43
Total Count of Early Prehistoric		1	0	1
Total Count of Late Archaic		30	14	44
Total Count of Middle Archaic		10	1	11
Total Count of Early Archaic		3	0	3

Woodlands, the Oak Woodlands-Prairie Riverine, and the Prairie habitats.

Conclusions

The GIS database of mortuary sites across the Gulf Coastal Plains shows the highest number of burials occurring in the Prairie during the Late Archaic and Late Prehistoric. Throughout the Late Archaic and Late Prehistoric, more sites contained multiple interments than single burials. The South Texas Brush Country is the only habitat to contain more single interment sites than multiple interment sites. All but one of the single interment sites dates to the Late Archaic or Late Prehistoric, while several multiple interment sites date earlier in the Oak Woodlands, Prairie-Riverine and Prairie habitats.

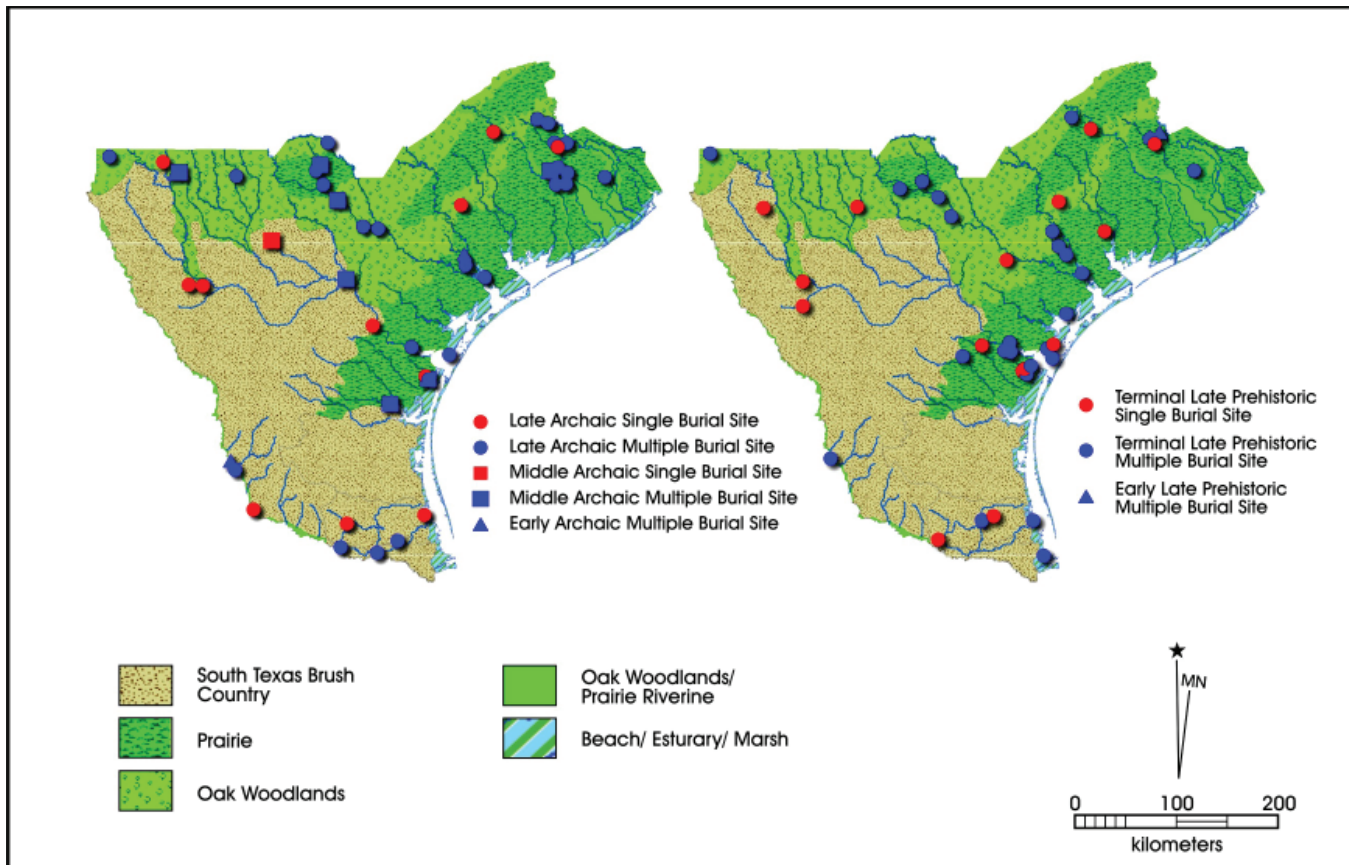


Figure 13-4. *Distribution of single and multiple interment sites across time and habitat.*

While the majority of the cemetery sites are located within a riverine setting, they are not necessarily clustered within the most productive habitats at the regional level. That is, if the placement of cemeteries was intended to mark territorial ownership of hyper-productive resources, we would expect that sites with human interments would be primarily found in the Prairie-Riverine habitat of the Coastal Plains and Marshes physiographic region. Instead, sites with human interments are distributed across a broad range of physiographic regions including the South Texas Brush Country, the Prairie, and the Oak Woodlands. This broad distribution shows that the locations of cemeteries are not exclusive to hyper-productive resource zones. Instead, they appear to be located in all

settings that are favorable to re-occupation and/or revisitation conditioned by a mix of resources that is typical of all riverine settings. Such a distributional pattern favors the interpretation that sites with human interments are located in “persistent places,” settings that are intermittently or regularly revisited and reoccupied sometimes continuously for generations while other times forgotten and reoccupied after centuries of abandonment. As such, the creation of cemetery sites as recognized by archeologists is more likely to represent a combination of short term processes that reflect cultural memory and identity and long term processes and conditions that favor reoccupation of landscape features conditioned by resource availability.

Chapter 14: Summary and Conclusions

Steve A. Tomka

The Cayo del Oso site (41NU2) has a long history of archeological work dating back to the early 1900's. A comprehensive summary of the early investigations at the site has been provided by the Center for Archaeological Research in the first of the two volume series published four years ago (Jackson et al. 2004). The present report represents the second and final volume in the series focused on archeological work conducted at the site.

This volume summarizes the archeological monitoring conducted in association with improvements made by TxDOT to Ennis Joslin Road and Ocean Drive in the vicinity of 41NU2. The archeological work was necessary to address the requirements of Section 106 of the National Historic Preservation Antiquities Code of 1966, as amended, the implementing regulations of 36 CFR Part 800, and the Texas Antiquities Code; because, the construction impacts land owned by the State of Texas and because the Federal Highways Administration (FHWA) is partially funding the project.

The archeological work carried out in support of the construction work consisted of monitoring of the excavation of trenches intended for utility line placement or relocation. The fieldwork began in August of 2000 and continued through February 2006. The purpose of the monitoring was to identify any cultural remains that may be within areas of impact and to "...complete the documentation of features and the exhumation of any human burials as quickly as possible to allow the construction contractors to regain access to the area." The work was conducted under WA #57311SA002, established between the University of Texas at San Antonio and the Texas Department of Transportation.

In the introductory chapters of this document we provide background to the history of the construction project (Chapter 1), the geology, climate and environmental setting of the site (Chapter 2) and briefly review the culture history of the project area (Chapter 3). In the next three chapters we focus on the monitoring activities carried out at 41NU2. In Chapter 4 we summarize the field methods and the monitoring protocol that was initially instituted and later modified as the project became more complicated due to the consistent finding of unmapped buried utility lines. In Chapter 5 we provide detailed descriptions of the monitoring efforts carried out during the project and each of the trenches that were inspected. We also describe the finding of six sets of human remains and of a relatively small number of artifacts that were collected during feature documentation episodes. The human remains consisted of one complete, intact burial

(Burial 2), three collections of disturbed, disarticulated human skeletal remains (Burials 3, 4, and 5), and two collections of partially disturbed human remains (Burials 6 and 7). In addition to these six, Ricklis (2004) recovered the remains of one more individual for a total of seven burials documented in association with the road improvements along Spur 3. The analysis and a complete description of these remains was completed by CAR under a separate report in compliance with the NAGPRA inventory of human remains from the site. This report was submitted to the TxDOT-ENV as part of a separate deliverable of this project. The small assemblage of artifacts contains worked shell, worked bone, two fragments of Native American ceramics (Rockport), and chipped stone artifacts. The two projectile points recovered are Late Archaic in age.

In Chapter 6 we discuss the results of a brief study of the relationship between paired sets of radiocarbon dates, one on carbonized wood and the other on shell. We conclude that marine shell can provide accurate, precise dates in the absence of other dating materials. This may especially be the case if samples of *Crassostrea virginica* are selected for dating. We also provided basic descriptive data on the 164 features as well as 11 occupation zones documented at the site during this monitoring project.

In the next three chapters we develop the theoretical framework (Chapter 7), describe the spatial and temporal structure of the resources (Chapter 8) and develop general and some specific expectations regarding the land use and subsistence strategies of prehistoric hunter-gatherers. In Chapter 7 we provide a detailed review of hunter-gatherer land use during the early 16th century as seen from the ethnohistoric record and describe the basics of human behavioral ecology that serves as the foundation of the theoretical models developed in Chapter 9. Chapter 8 is an exhaustive summary of regional resource structure and Chapter 9 develops some general expectations based on the habitat choice model and identifies some specific expectations regarding land use and subsistence at 41NU2. In Chapter 9 we also develop some research problems and proposed avenues of research to guide the analytical approaches that were followed in the remaining chapters of the report. The next four chapters contain the results of site-specific and regional-scale analyses, including paleoclimate (Chapter 10), subsistence (Chapter 11), land use (Chapter 12), and mortuary practices (Chapter 13).

Several distinct data types are brought together to investigate trends in paleoclimate including pollen,

phytoliths, macrobotanical remains, diatoms, ostracods, and geomorphological data. The results of these specialized analyses are provided in individual appendixes (A-F). The paleoclimatic investigations center of soil columns obtained from three locations at or near the site. It is difficult to find correspondence between the various data sets across all three locations. Nonetheless, the data seem to indicate some broad changes in climate over time from relatively mesic conditions and mesic-adapted vegetation communities to more xeric conditions over time, with a brief shift back to more mesic conditions in the last 1000 years. The dune cores (Cores 4 and 6) document very localized geomorphic processes that make the patterns they reflect more indicative of small-scale processes rather than regional trends. The analysis of the pollen from Core 2 by two different laboratories resulted in dramatically different conclusions emphasizing the need to further investigate the nature of the differences to ensure that the results of future analyses are reliable.

Chapter 11 reports on the analysis of the faunal remains recovered from heavy fractions derived from the floatation of feature fill. The analysis of the species of fish present in the deposits suggests a late winter/early spring season of acquisition and site use. The splintered nature of the deer-sized faunal remains may be indicative of bone grease processing or some degree of nutritional stress experienced by the site's inhabitants. Numerous lenses of fish bone across different parts of the dunes indicate that small to medium fish were commonly procured and perhaps minimally processed. The comparison of subsistence practices at the regional level indicates that mammals, birds, fish and molluscs have been incorporated in the diet of hunter-gatherers living in the Grass Flats, the Bay-Fluvial and Prairie/Riverine habitats. Nonetheless, the data also suggests that the hunter-gatherers living in each of these habitats tended to emphasize the exploitation of resources found within them. This pattern is supported by the results of stable carbon and nitrogen isotope analyses of coastal, riverine, and inland populations (Johnson and Hard 2008) and also matches the available ethnohistoric information (Ricklis 1996).

Chapter 12 indicates that regional land use was very much centered on prairie-riverine habitats and the earliest occupations across the region occurred in this habitat. Site sizes, on the average, tend to be larger in the riverine setting than in any other habitat and small sites are under-represented in this setting. In the beach-fluvial habitat sites tend to be dominated by small sites but the mean site size increases on the beach most likely due to the merging of palimpsest occupations. A gradual reorientation in the use of coastal habitats is noted in the Late Archaic and this pattern changes to a significant shift to the exploitation of coastal aquatic resources during the Late Prehistoric Period. This change

may have been prompted by population packing within the Prairie/Riverine habitat and may be a reflection of hunter-gatherer intensification of aquatic resource extraction.

The analysis of hunter-gatherer mortuary practices through sites containing human remains shows that the bulk of the sites with human remains are found in the Oak Woodlands/Prairie Riverine habitat. The linear arrangement of cemetery sites suggests that hunter-gatherer territories, if present, centered on stream valleys and drainages. The temporal patterning of the sites with human remains does suggest that the increase in sites with human interments corresponds to increases in population densities. However, the fact that cemetery sites are found in riverine settings throughout the broad South Texas Plains, the Oak Woodlands and Prairie does not conform to the expectation that they should be primarily limited to hyper-productive habitats if they are intended as territorial signals. The occurrence of cemetery sites in riverine settings cutting across physiographic regions characterized by low productivity suggests that cemetery formation is a product of the repeated reuse of selected places across the landscape. Overall, the distribution of cemetery sites suggests that they are created in settings that are revisited and reoccupied repeatedly either because they contain resources, facilities or other features that draw populations to them over long periods of time (generations).

Overall, the lengthy project produced little data compared to other archeological projects of comparable duration. However, this is understandable given the monitoring focus of the investigations. The site-specific paleoenvironmental results are encouraging although somewhat troubling. The regional-scale analyses produced a mixed bag of results with some very intriguing patterns that support expectations derived from the model and other that simply did not produce sufficient data to adequately test and investigate its implications.

It is possible that different monitoring protocols could have produced more systematic and larger amounts of data. Construction monitoring was driven by the knowledge that human burials may be found anywhere within the project limits. Therefore, archeological monitoring was a necessity. However, and as amply evident in the report, limiting data collection to the context of monitoring has reduced the quantity, quality and variety of data types that could be obtained. During the early phases of the project longer stoppages in construction activities allowed for more complete sampling of features but these stoppages understandably tended to interfere with the pace of the construction. This is the classical dilemma of "rescue archeology" within the context of Cultural Resource Management. Features are

excavated and documented as they are encountered and there is a premium on time since construction crews and equipment remain idle while archeologists are at work.

One approach that could have been used in conjunction with or perhaps in advance of construction monitoring would have been to design an overall project schedule that allowed time for more intensive investigation of material-rich portions of the Right-of-Way well in advance of the construction activities. Hand-excavations in portions of the ROW not scheduled for immediate construction impact would have allowed detailed investigations of some areas that would have yielded representative and statistically significant amounts of data for comparative analyses while at the same time not interfering with construction schedules. Perhaps such an approach would have then allowed for monitoring to focus on potential human remains and no attempts to collect data on features and occupation zones that would have been adequately sampled during the pre-construction phase of the archeological investigations. To implement such a strategy would have required extensive planning well in advance of the construction and the availability of personnel for both monitoring and data-recovery activities.

One of the glaring weaknesses of the report is that while we hoped to quantitatively test some expectations derived from the Habitat Choice Model through cost-benefit analysis, early on it became evident that the necessary data is simply not available in the State. Quantitative tests of some of the premises of Optimal Foraging Models have proven to be very instructive and have yielded new insights into the parameters and expectations derived from them (see Bettinger et al. 2006; Ugan 2005; Zeanah 2000). They clearly provide worthwhile avenues of future research and if archeologists in the State are to take full advantage of them, the systematic collection of appropriate data (i.e., caloric yield data) combined with experimental cost-benefit analysis (i.e., quantifying the energetic costs of certain extractive activities) is greatly desired. Similarly, detailed understanding of the structure of the economically significant resources available to prehistoric hunter-gatherers is also lacking. And, while modern resource structure does not necessarily reflect prehistoric resource availability and distribution, we need to continue to strive to understand how resources may have changed over time to be able to construct more realistic models for understanding

hunter-gatherer behavior. To this end, we need to continue to search for and collect data that will allow us to reconstruct in increasingly greater detail regionally specific climatic sequences that in turn will allow us to more precisely approximate past resource structure.

In addition to collecting new data types, it is also important to present data from ongoing and new testing and data recovery projects in a consistent and systematic manner. The study of regional-scale phenomena is hampered at this time by the difficulty of extracting the desired data from technical reports. When one is finally able to locate it, it is often not analyzed in a consistent manner or not presented in sufficient detail. Some degree of standardization of research and analytical approaches is needed to allow researchers the opportunity to take advantage of the growing set of data being collected with each cultural resource management testing and data recovery project. While standardization cannot stifle creativity and openness to trying new research avenues, it is necessary to allow archeologists to investigate hunter-gatherer adaptations and behavior at the regional-scale.

The Texas Archeological Sites Atlas is a good beginning towards a repository of data associated with projects and individual sites. However, the data is not research-friendly and often those using the information are left to print out site forms and extract relevant site and project characteristics by hand to build usable databases. Similarly, many of the testing and data recovery reports are available only in a limited manner (the THC library), and comprehensiveness of one's regional comparisons often depend on how complete the company's research library is. It is hoped that this will be rectified as the Texas Historical Commission moves to make pdf copies of all reports available.

Having outlined in general terms some of the weaknesses of this report and also having described the data gaps that have influenced the quality of the product, it is important to indicate that these omissions are not provided as excuses for the quality of the final product. Archeological knowledge is to a large extent cumulative and it is hoped that researchers using this report can build on the current state of knowledge as we have attempted to add to the research of others who are interested in coastal archeology.

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APPENDIX A:
Pollen and Phytolith Analysis of Stratigraphic Core Samples
from the Cayo Del Oso Site (41NU2), Copus Christi, Texas

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Introduction

Pollen and phytolith analysis was conducted on three stratigraphic core samples from the Cayo del Oso site (41NU2). The site is located along the northwestern shore of Oso Bay, Corpus Christi, Texas. Previous excavations have indicated that the site was a prehistoric campsite and cemetery. Radiocarbon dates from the test units and the burial ranged from 4400 B.P. to 1300 B.P. Core samples submitted for analysis were derived from pond sediments (Core 2) and dune deposits (Cores 4 and 6). Pollen and phytolith analysis of these samples was undertaken to provide paleoenvironmental reconstruction data for the Oso Bay Site and coastal South Texas.

Methods

Pollen

A chemical extraction technique based on flotation is the standard preparation technique used in this laboratory for the removal of the pollen from the large volume of sand, silt, and clay with which they are mixed. This particular process was developed for extraction of pollen from soils where preservation has been less than ideal and pollen density is lower than in peat.

Hydrochloric acid (10%) is used to remove calcium carbonates present in the soil, after which the samples are screened through 150 micron mesh. The samples are rinsed until neutral by adding water, letting the samples stand for 2 hours, then pouring off the supernatant. A small quantity of sodium hexametaphosphate is added to each sample once it reaches neutrality, then the samples are allowed to settle according to Stoke's Law in settling columns. This process is repeated with ethylenediaminetetraacetic acid (EDTA). These steps remove clay prior to heavy liquid separation. The samples are then freeze dried under a partial vacuum. Sodium polytungstate (SPT), with a density 1.8, is used for the flotation process. The samples are mixed with SPT and centrifuged at 1500 rpm for 10 minutes to separate organic from inorganic remains. The supernatant containing pollen and organic remains is decanted. Sodium polytungstate is again added to the inorganic fraction to repeat the separation process. The supernatant is decanted into the same tube as the supernatant from the first separation. This supernatant is then centrifuged at 1500 rpm for 10 minutes to allow any silica remaining to be separated from the organics. Following this, the supernatant is decanted into a 50 ml conical tube and diluted with distilled water. These samples are centrifuged at 3000 rpm to concentrate the organic fraction in the bottom of the tube. After rinsing the pollen-

rich organic fraction obtained by this separation, all samples receive a short (20-30 minute) treatment in hot hydrofluoric acid to remove any remaining inorganic particles. The samples are then acetolated for 3-5 minutes to remove any extraneous organic matter.

A light microscope is used to count the pollen to a total of approximately 200 pollen grains at a magnification of 500x. Pollen preservation in these samples varied from good to poor. Comparative reference material collected at the Intermountain Herbarium at Utah State University and the University of Colorado Herbarium was used to identify the pollen to the family, genus, and species level, where possible.

Pollen aggregates were recorded during identification of the pollen. Aggregates are clumps of a single type of pollen and may be interpreted to represent pollen dispersal over short distances or the introduction of portions of the plant represented into an archeological setting. Aggregates were included in the pollen counts as single grains, as is customary. The presence of aggregates is noted by an "A" next to the pollen frequency on the pollen diagram. Pollen diagrams are produced using Tilia, which was developed by Dr. Eric Grimm of the Illinois State Museum. Total pollen concentrations are calculated in Tilia using the quantity of sample processed in cubic centimeters (cc), the quantity of exotics (spores) added to the sample, the quantity of exotics counted, and the total pollen counted and expressed as pollen per cc of sediment.

Indeterminate pollen includes pollen grains that are folded, mutilated, and otherwise distorted beyond recognition. These grains are included in the total pollen count since they are part of the pollen record. The microscopic charcoal frequency registers the relationship between pollen and charcoal. The total number of microscopic charcoal fragments was divided by the pollen sum, resulting in a charcoal frequency that reflects the quantity of microscopic charcoal fragments observed, normalized per 100 pollen grains.

Pollen analysis also includes examination for and identification of starch granules to general categories, if they are present. Starch granules are a plant's mechanism for storing carbohydrates. Starches are found in numerous seeds, as well as in starchy roots and tubers. The primary categories of starches include the following: with or without visible hila, hilum centric or eccentric, hila patterns (dot, cracked, elongated), and shape of starch (angular, ellipse, circular, eccentric). Some of these starch categories are typical of specific plants, while others are more common and tend to occur in many different types of plants.

Phytoliths

A pressurized (closed) microwave digestion technique was used to oxidize the majority of the organic fraction of the sediment samples. Specific protocols were developed by PRI that were based on a few previous studies (Parr, et al. 2001; Parr, et al. 2004). One gram of sediment was placed in a glass vial with 4 ml nitric acid (HNO₃), 2 ml hydrochloric acid (HCl), and 1 ml of hydrogen peroxide (H₂O₂). Each glass vial was placed and sealed into a Teflon vessel with outer casement and screw cap seal. One vessel was connected to a pressure sensor transducer for real-time pressure monitoring by the in-built microwave computer system (Floyd Inc. Remote Microwave System 150). Samples were programmed for exposure to 40 psi at a dwell time of 60 minutes, using 100% magnetron power, then allowed to cool and depressurize. Samples were then removed from the microwave and rinsed to neutral in 50 ml centrifuge tubes (4 X 3000 rpm for 2 min). One ml of ethylenediaminetetraacetic acid (EDTA) was added to each sample, which was then placed in an ultrasonic bath and sonicated for approximately three minutes. EDTA aids in the removal of trace metals and humic substances undigested by acid oxidation. Samples were then rinsed to neutral.

Because sediment deposits from the Oso Bay Site contained a large fraction of silt-sized particles, many of which are similar in size to our target phytoliths, standard PRI phytolith methodology had to be customized. After the EDTA treatment, samples were sieved through a 50 µm mesh into 50 ml centrifuge tubes, rinsed to neutral (4 X 3000 rpm for 2 min), transferred to 15 ml centrifuge tubes, and then dried using a partial vacuum drying system (lyophilization). The lyophilized samples were then subjected to heavy liquid flotation to separate silica phytoliths from non-silica particles. Separation trials were run using sodium polytungstate densities of 2.3, 2.2 and 2.1 and Cadmium iodide at densities of 2.2 and 2.15. All samples were then treated with cadmium iodide at a density of 2.15, as this was found to provide the best separation of phytoliths from the silt fraction with very little loss of phytoliths to the heavy fraction. Samples were mixed with cadmium iodide and centrifuged at 1500 rpm for 10 minutes. The supernatant containing phytoliths was decanted. Cadmium iodide was again added to the heavy fraction to repeat the separation process. The supernatant was decanted into the same tube as the supernatant from the first separation. This supernatant was then centrifuged at 1500 rpm for 10 minutes to allow any non-phytolith silica and other particles remaining to be separated from the phytoliths. Following this, the supernatant was decanted into a new centrifuge tube, then diluted with distilled water and centrifuged at 3000 rpm for 10 minutes to concentrate the phytolith fraction in the bottom of the tube. Additional rinses were then conducted to wash any remaining heavy liquid

from the sample. Because of the initial small sample size started, most of the clay-sized particle fraction was removed during previous rinse and centrifugation steps. If necessary, the remaining clay-sized particles were removed with short-time centrifugation in 15 ml tubes (typically 1 or 2 spins at 3000 RPM for 1.5 min). After several alcohol rinses, samples were transferred to storage vials, and a small portion mounted in Cargille Type-A immersion oil (refractive index 1.515), with two opposite coverslip corners affixed with nail polish to allow for controlled phytolith rotation and counting with a light microscope at a magnification of 500x.

Phytolith Review

Phytoliths are silica bodies produced by plants when soluble silica in the ground water is absorbed by the roots and carried up to the plant via the vascular system. Evaporation and metabolism of this water result in precipitation of the silica in and around the cellular walls. Opal phytoliths, which are distinct and decay-resistant plant remains, are deposited in the soil as the plant or plant parts die and break down. They are, however, subject to mechanical breakage and erosion and deterioration in high pH soils. Phytoliths are usually introduced directly into the soils in which the plants decay. Transportation of phytoliths occurs primarily by animal consumption, gathering of plants by humans, or by erosion or transportation of the soil by wind, water, or ice.

The three major types of grass short-cell phytoliths include festucoid, chloridoid, and panicoid. Smooth elongate phytoliths are of no aid in interpreting either paleoenvironmental conditions or the subsistence record because they are produced by all grasses. Phytoliths tabulated to represent “total phytoliths” include the grass short-cells, bulbiform, trichome, elongate, and dicot forms. Frequencies for all other bodies recovered are calculated by dividing the number of each type recovered by the “total phytoliths”.

The festucoid class of phytoliths is ascribed primarily to the Subfamily Pooideae and occur most abundantly in cool, moist climates. However, Brown (1984) notes that festucoid phytoliths are produced in small quantity by nearly all grasses (mostly rondel-type phytoliths). Therefore, while they are typical phytoliths produced by the Subfamily Pooideae, they are not exclusive to this subfamily. Chloridoid phytoliths (short saddles) are found primarily in the Subfamily Chloridoideae, a warm-season grass that grows in arid to semi-arid areas and require less available soil moisture. Chloridoid grasses are the most abundant in the American Southwest (Gould and Shaw 1983:120). Bilobates and polylobates (lobates) are produced mainly by panicoid grasses, although a few of the festucoid grasses also produce these forms. Panicoid

phytoliths occur in warm-season or tall grasses that frequently thrive in humid conditions. Twiss (1987:181) also notes that some members of the Subfamily Chloridoideae produce both bilobate (Panicoid) and Festucoid phytoliths. "According to (Gould and Shaw 1983:110) more than 97% of the native US grass species (1,026 or 1,053) are divided equally among three subfamilies Pooideae, Chloridoideae, and Panicoideae" (Twiss 1987:181).

Buliform phytoliths are produced by grasses in response to wet conditions and are to be expected in wet habitats of floodplains and other places. Trichomes represent silicified hairs, which may occur on the stems, leaves, and the glumes or bran surrounding grass seeds.

Conifers produce opal silica phytoliths in their inner bark and needles. Polyhedral phytoliths are noted to be observed in leaves (Bozarth 1993), while the authors have observed the blocky forms in bark reference samples.

A rapidly growing body of phytolith studies, many of which using three-dimensional typological descriptions, are allowing a much finer resolution of taxa identification, often to the subfamily and genus level. This is especially true with the grass family, where diagnostic forms are now accepted for many additional subfamilies and genera such as: Bambusoideae (bamboo), Ehrhartoideae (rice), tribe Oryzeae (*Oryza*, *Zizania*), Aristidoideae (threeawn, needlegrass), Arundinoideae (reeds), and *Zea mays* (maize, corn) (Piperno 2006). Determining local and regional levels of diagnostic morphotypes can also greatly improve taxa identification.

Lobate Phytolith Classification

Study areas, such as this one, with an abundance of lobate phytoliths (e.g., bilobates, polylobates) allow for the use of a detailed lobate phytolith classification. Lobate phytoliths are most characteristic of the grass subfamily Panicoideae (e.g., warm and moist season tall grass prairie species, sugarcane, millet), but are also sometimes found in the Chloridoideae (e.g., warm and dry season short grass prairie species, cord grass), Ehrhartoideae (e.g., white rice, North American wild rice) and Bambusoideae (bamboos). Lobate phytoliths are typically comprised of two lobes joined by a shank, and have been described as resembling a dumbbell. The lobate phytolith classification used here is based on: (1) the length of the lobate shank and (2) the shape of the outer margin of the two lobes. This lobate classification follows that used by Lu and Lu (2003) and their analysis of 250 modern grass species. Their work demonstrated a link between lobate phytolith morphology, grass taxonomy and environmental conditions. Detailed bilobate phytolith classification can potentially

detect the presence of economically and culturally significant grass species. In addition, detailed bilobate classification is more sensitive to subtle ecological and precipitation change than that detectable by traditional festucoid, chloridoid and panicoid relative abundance. Bilobate classification graphs of relative abundance are summed independently, but are based on the overall phytolith count (typically 300 grains).

Other Siliceous Microfossils

Diatoms and sponge spicules also were noted. Pennate diatoms are cosmopolitan, occurring in many sediments, and they indicate at least some soil moisture. Sponge spicules represent fresh water sponges. Diatoms are single-celled algae with a siliceous cell wall. They grow in a wide range of aerophilous habitats, including on wet plants and rocks, in damp soils, marshes, wetlands, mudflats, and all types of standing and flowing aquatic habitats. Their silica cells often are preserved in sedimentary deposits. Because individual taxa have specific requirements and preferences with respect to water chemistry, hydrologic conditions, and substrate characteristics, the presence (and subsequent identification to the species level) of diatoms in paleoenvironmental context can provide information about the nature of the local environment. These data, coupled with input about local geology, hydrology, soil characteristics, pollen and phytoliths, provide evidence of the paleoenvironmental setting. In the context of phytolith samples, diatoms are noted but not identified beyond the split of "pennate" and "centric" forms. Centric diatoms often indicate wet conditions, while at least some of the pennate diatoms are cosmopolitan, occurring nearly everywhere. Both diatoms and sponge spicules can be transported with sediment. As an illustration, recovery of sponge spicules in upland soils is noted to accompany loess deposits derived from floodplains in Illinois (Jones and Beavers 1963).

Discussion

Pollen and phytolith analysis was conducted on three stratigraphic core samples (Table A-1) from the Cayo Del Oso site (41NU2). The site is located along the northwestern shore of Oso Bay, Corpus Christi, Texas. The modern climate is considered subtropical and semi-arid. Mean annual precipitation for the area averages 74 cm per year. Mean annual evaporation rates average from 90 to 115 cm per year, but may range as high as 150 cm per year during times of drought. Southeasterly prevailing winds are characteristic of the Texas Gulf coast for most of the year. Strong northerly frontal passages are common in winter (Nicolau 2001). An ecological gradient exists along the Texas coast, with higher rainfall amounts and greater freshwater inflows

Table A-1. Provenience Data for Samples from the Cayo del Oso Site (41NU2)

Sample No.	Core No.	Segment	Zone	Depth within segment	Description	Analysis
103	2	1	8	97-101		Pollen & Phytolith
104	2	2	8	8-13		Pollen & Phytolith
105	2	2	9	18-23		Pollen & Phytolith
106	2	2	9	28-32		Pollen & Phytolith
107	2	2	9	37-41		Pollen & Phytolith
108	2	2	10	46-51		Pollen & Phytolith
109	2	2	10	56-61		Pollen & Phytolith
110	2	2	10	65-70		Pollen & Phytolith
111	2	2	10	74-79		Pollen & Phytolith
112	2	2	10	84-89		Pollen & Phytolith
113	2	2	10	96-101		Pollen & Phytolith
114	2	3	10	28-32		Pollen & Phytolith
115	2	3	10	35-40		Pollen & Phytolith
116	2	3	11	44-48		Pollen & Phytolith
117	2	3	11	53-57		Pollen & Phytolith
182	4	1	5	40-43		Phytolith
183	4	1	6	46-49		Phytolith
184	4	1	8	57-60		Phytolith
185	4	1	9	68-70		Phytolith
186	4	1	11	77-80		Phytolith
187	4	1	12	85-87		Phytolith
188	4	1	13	90-93		Phytolith
189	4	2	14	0-3		Phytolith
190	4	2	16	12-14		Phytolith
191	4	2	17	20-23		Phytolith
192	4	2	17	30-33		Phytolith
193	4	2	18	35-38		Phytolith

Table A-1. Continued...

Sample No.	Core No.	Segment	Zone	Depth within segment	Description	Analysis
194	4	2	19	46-49		Phytolith
195	4	2	20	53-56		Phytolith
196	4	2	21	56-59		Phytolith
197	4	2	22	59-63		Phytolith
198	4	2	23	70-73		Phytolith
199	4	2	23	80-83		Phytolith
200	4	2	24	92-95		Phytolith
201	4	3	25	7-10		Phytolith
202	4	3	26	14-17		Phytolith
203	4	3	26	25-28		Phytolith
204	4	3	27	35-38		Phytolith
279	6	1	4	28-30		Phytolith
280	6	1	5	32-34		Phytolith
281	6	1	6	40-42		Phytolith
282	6	1	7	47-49		Phytolith
283	6	1	8	53-55		Phytolith
284	6	1	9	70-73		Phytolith
285	6	1	9	80-83		Phytolith
286	6	1	9	90-93		Phytolith
287	6	1	10	53-56		Phytolith
288	6	2	11	75-78		Phytolith
289	6	2	18	88-90		Phytolith
290	6	3	13	5-8		Phytolith
291	6	3	14	25-28		Phytolith
292	6	3	15	42-45		Phytolith
293	6	3	16	59-62		Phytolith
294	6	3	17	70-73		Phytolith

in the northern half of the coastal region giving rise to the establishment of salt marshes dominated by smooth cordgrass (*Spartina alterniflora*, subfamily Chloridoideae). In the south, low annual rainfall and very little freshwater inflow result in the creation of vast areas of algal flats (wind-tidal flats) with little to no emergent salt marsh vegetation. The Corpus Christi Bay Complex, positioned just south of the midway point on the Texas Gulf coast, serves as the ecotone between these two distinct habitat types. Stands of *Spartina alterniflora* become less common from Corpus Christi Bay south, and are replaced by algal flats, which become larger and more common. Between 9,000 and 3,000 B.P. sea level rose unevenly (stable, slow and rapid), after approximately 3,000 B.P. sea level has been at about its modern level (Ricklis n.d.). Previous archeological excavations from the 1920s have indicated that the site was a prehistoric campsite and cemetery. Radiocarbon dates from test units and a burial ranged from 4,400 B.P. to 1,300 B.P. Construction activities associated with road improvements have required recent and ongoing excavations and monitoring to mitigate damage to the site. Since 2000, monitoring and excavations have yielded intact stratified cultural deposits and multiple features and burials. In addition, an area of pond sediments and two dune deposits with up to 12 identifiable paleosols have been identified. Core samples submitted for pollen and phytolith analysis are derived from pond sediments (Core 2) and dune deposits (Cores 4 and 6). Pollen and phytolith analysis of these samples should provide paleoenvironmental reconstruction data for the Oso Bay Site and coastal South Texas.

Phytolith interpretations from south and southeastern coastal sediments of North America have been greatly enhanced by the work of Lu and Liu (2003; 2003; 2005). They have expanded our knowledge of phytolith production in many common coastal plants, especially within the grasses (Poaceae), and have associated modern phytolith assemblages with major coastal vegetation zones (salt marsh, brackish/fresh marsh, swamps, forests, interdune meadows, sand dunes). They have also developed a bilobate phytolith classification that can be used to detect environmental change. This classification was employed here to detect changes in grass species associated with forest edge / scrubland ecotones, sandy soils and shallow water/shoreline margins. This was made possible here after matching bilobate morphotypes and their associated grass species with habitat preference information.

Core 2 (pond sediments)

Core 2 was collected in pond sediments to interpret paleoenvironmental conditions in this area by examining several different data sets. Pollen and phytoliths were extracted from a set of stratigraphic samples from this core

to obtain information concerning the local and regional vegetation communities.

Plants considered typical along the northern Gulf Coast (including Texas) include beach pioneers and plants growing in humid and “wet slack” areas. Beach plants of this group (excluding grasses) are represented in the pollen record as: Brassicaceae (perhaps *Cakile* sp.), Onagraceae (perhaps *Oenothera*), and *Sesuvium*. Plants (excluding grasses) that grow in humid slacks and on beaches where the ground water table is high include Cyperaceae, *Hydrocotyle bonariensis* (Apiaceae), and *Salicornia bigelovii* (a Chen-am). Recovery of pollen from these families does not necessarily indicate that these specific genera or species are represented. Humic slacks support grasses, sedges, and *Salicornia bigelovii* (a Chen-am). Grasslands and thickets include grasses, prickly pear cactus (*Opuntia*), and a variety of plants not represented in the pollen record. The wide variety of grasses growing in the area limit the use of Poaceae pollen to interpret the vegetation community represented. Instead, the strength of the phytolith will be used to interpret the presence of specific groups of grasses (Moreno-Casasola 1988).

The pollen record exhibits Poaceae, Chen-am, and High-spine Asteraceae pollen (Figure A-1, Table A-2) as the three co-dominants, representing grasses, various members of the goosefoot families and perhaps amaranth that include at least *Salicornia bielovii*, which is noted to grow in the area today in humid slacks and on beaches. The High-spine Asteraceae is a large group of plants including most of the sunflower family that is not accounted for in the Liguliflorae (chicory tribe) or Low-spine Asteraceae (ragweed, cocklebur, bursage, and marshelder).

The pollen and phytolith records are both divided into matching zones for ease of discussion of the interpretations. General trends in the pollen record include large changes in the microscopic charcoal record (Figure A-1). Quantities of microscopic pieces of charcoal, which are calculated relative to the quantity of pollen in each sample, are moderately small in Zone E, increasing in Zone D, and generally higher in Zone C, declining in Zone B, then increasing in Zone A, with a decline noted in the uppermost sample. From the microscopic charcoal record alone, it is not possible to identify whether these fluctuations correspond with human activity and represent short distance wind transport as ash released from hearths or whether they represent longer distance transport from local or regional wild fires.

The phytolith record indicates that there has been some subtle, long-term environmental change as well as a few punctuated events and/or short term changes strong

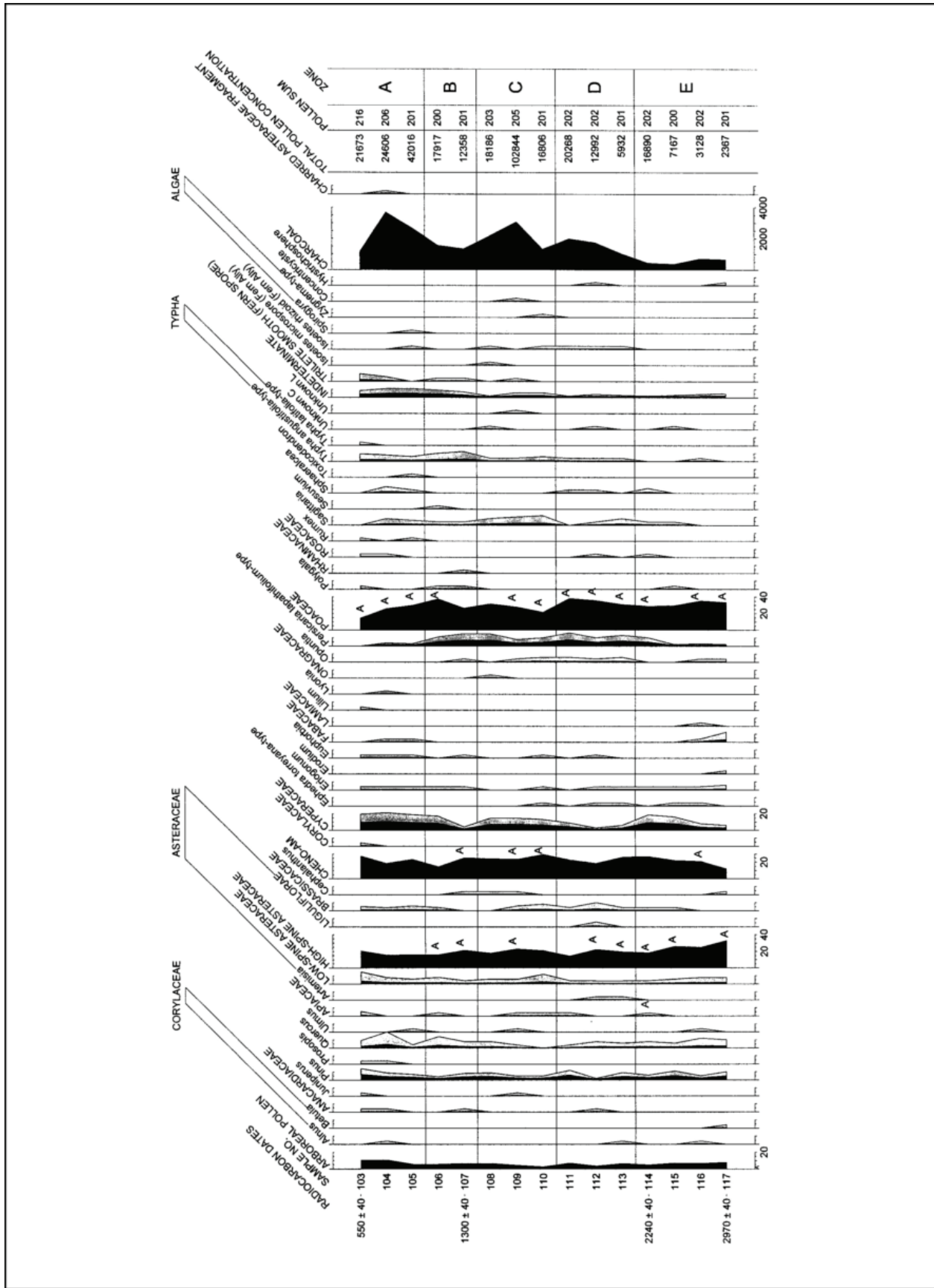


Figure A-1. Pollen diagram, Core 2, Site 41NU2, Cayo del Oso, Texas.

Table A-2. Pollen Types Observed in Samples from the Cayo del Oso Site (41NU2).

Scientific Name	Common Name
ARBOREAL POLLEN:	
Anacardiaceae	Sumac family
Corylaceae:	Birch family
<i>Alnus</i>	Alder
<i>Betula</i>	Birch
<i>Juniperus</i>	Juniper
<i>Pinus</i>	Pine
<i>Prosopis</i>	Mesquite
<i>Quercus</i>	Oak
<i>Ulmus</i>	Elm
NON-ARBOREAL POLLEN:	
Asteraceae:	Sunflower family
<i>Artemisia</i>	Sagebrush
Low-spine	Includes ragweed, cocklebur, sumpweed
High-spine	Includes aster, rabbitbrush, snakeweed, sunflower, etc.
Liguliflorae	Chicory tribe, includes dandelion and chicory
Brassicaceae	Mustard family
<i>Cephalanthus</i>	Buttonbush
Cheno-am	Includes the goosefoot family and amaranth
Corylaceae	Hazel family
Cyperaceae	Sedge family
<i>Ephedra torreyana</i> -type (includes <i>E. torreyana</i> , <i>E. trifurca</i> , and <i>E. antisyphilitica</i>)	Ephedra, Jointfir, Mormon tea
<i>Eriogonum</i>	Wild buckwheat
<i>Erodium</i>	Storksbill, Heron-bill, Filaree
Fabaceae:	Bean or Legume family
Lamiaceae	Mint family
<i>Lilium</i>	Lilley
<i>Lyonia</i>	staggerbush

Table A-2. Continued...

Scientific Name	Common Name
Onagraceae (Gaura)	Evening primrose family
<i>Opuntia</i>	Prickly pear cactus
<i>Persicaria-tapanthifolium</i> -type	Knotweed, Spotted lady's thumb
Poaceae	Grass family
<i>Polygala</i>	Milkwort
Rhamnaceae	Buckthorn family
Rosaceae:	Rose family
<i>Rumex</i>	Dock
<i>Sagittaria</i>	Arrowweed
<i>Sesuvium</i>	Sea purslane
<i>Sphaeralcea</i>	Globe mallow
<i>Toxicodendron</i>	Poison ivy
<i>Typha angustifolia</i> - type	Cattail
<i>Typha latifolia</i> - type	Cattail
Indeterminate	Too badly deteriorated to identify
SPORES:	
Trilete Smooth	Fern spore
Isoetes Microspores	Fern Ally
Isoetes Rhizoid	Fern Ally
ALGAE:	
<i>Spirogyra</i>	Algal body representing a filamentous green algae common in freshwater habitats
<i>Zygnema</i> -type	Algal body
OTHER:	
Concentricyst	Fungal spore, indicator of wet, oxidized conditions
Hystrichosphere	Geologic plankton
Charcoal	Microscopic charcoal fragment
Charred Asteraceae tissue fragment	Charred tissue fragment from a member of the sunflower family
Total pollen concentration	Quantity of pollen per cubic centimeter (cc) of sediment.

enough to be detectable with the sampling interval strategy employed (Figure A-2). The large sampling intervals, both within and between samples, limits the availability to detect punctuated events and has an overall smoothing effect on the environmental record. Despite being limited to a course-scale, the bilobate phytolith classification allows us additional interpretive data, and likewise indicates both long-term trends and short-term events (Figure A-3). Conjoining samples exhibiting similar phytolith assemblages have been grouped together into zones. Discussion will start with the chronologically lowest zone.

Radiocarbon dates of 2970 ± 40 RCYBP and 2240 ± 40 RCYBP bracket Zone E at the base of the stratigraphic record. This interval exhibits relatively stable Poaceae pollen frequencies and declining High-spine Asteraceae and increasing Chenom and Cyperaceae pollen frequencies. Small quantities of *Alnus* and *Betula* pollen noted low in this unit indicate the regional presence of alder and birch. Recovery of small quantities of *Pinus* and *Quercus* pollen throughout this zone and the remainder of the pollen record indicate the pines and oak grew as part of the regional vegetation throughout the past approximately 3,000 years. *Ulmus* pollen was observed once in this zone. Low-spine Asteraceae, Brassicaceae, *Cephalanthus*, *Eriogonum*, *Erodium*, Fabaceae, Lamiaceae, *Opuntia*, *Polygala*, Rosaceae, and *Sphaeralcea* pollen are noted in small quantities, indicating local growth of marshelder and similar plants, members of the mustard family, buttonbush, wild buckwheat, filaree, legumes, members of the mint family, prickly pear cactus, milkwort, members of the rose family, and globe mallow grew in the area. Increasing quantities of *Persicaria lapathifolium*-type and *Smilax* pollen, and the intermittent presence of *Typha* pollen represents a knotweed (pale persicaria or a similar plant) that prefers mesic habitats, greenbriar, which typically grows at the margins of woodlands, and cattails, which grows primarily in freshwater marsh habitats. Recovery of a small quantity of *Ephedra torreyana* pollen in this and succeeding levels of this record probably represents long distance wind transport from populations growing in drier habitats.

Samples from Zone E exhibit a phytolith assemblage that suggests a coastal freshwater marsh/pond habitat existed, as evidenced by the high abundance of bilobates (called dumbbells by Lu and Liu, 2005), smooth elongates, echinate elongates (called sinuate elongates by Lu and Liu, 2005), and buliforms (called squares and rectangles by Lu and Liu, 2005). Tribe Oryzae morphotypes are also relatively abundant in this zone and are likely derived from aquatic/wetland grasses *Zizaniopsis miliacea* (giant cutgrass) and *Leersia* spp. (Indian cutgrass). This zone also exhibited the highest relative abundance of pennate diatoms, usually associated with freshwater habitats of lakes and streams. Bilobate

indicators of forest (scrub) edge and understory grasses are in very low abundance, with bilobate indicators of sandy soil (beach and dune) grasses and shoreline and emergent aquatic grasses fairly high in abundance (Figure A-3). Beach and dune grass phytoliths are likely derived from *Panicum amarum*, *Uniola paniculata*, *Aristida desmantha*, *Cenchrus* spp. and *Spartina* spp. Phytoliths from wetland grasses are likely derived from *Leersia* spp. and *Zizaniopsis miliacea*. Thus, Zone E likely represents an expansive freshwater marsh habitat with dunes and tidal mudflats extending further than they do today because of slightly lower sea levels. In fact, at approximately 3,000 B.P., sea levels may have dipped slightly (≈ -1.0 m) from their long-term slowly rising trend since the Last Glacial Maximum (Flemming, et al. 1998; Milne, et al. 2005).

Zone D is marked by a continuation of the decline in High-spine Asteraceae pollen. Quantities of Chenom and Poaceae pollen fluctuate slightly in this zone. Cyperaceae pollen declines, then rises toward the top of this zone. Both *Opuntia* and *Persicaria lapathifolium*-type pollen increase and *Smilax* pollen retains its presence throughout most of this zone. *Typha* pollen is present in every sample. Pollen recovered from this zone exhibits indicators of beach and grasslands and thickets through the presence of *Opuntia* pollen. *Persicaria* grows along disturbed margins of wetlands, lakes and rivers. Its presence suggests periodic water level change and/or the presence of enough open water to create wave action, resulting in a margin of disturbance. The viney growth habit of *Smilax* and its current distribution in East Texas and the southeastern United States suggests wetter conditions, i.e. a positive water balance, than what exists today. *Typha* thrives in freshwater marshes. *Isoetes* (quillwort) rhizoids make their first appearance in Zone D. Nearly 150 species of quillworts are known. They grow in conditions from aquatic to semi-aquatic to drying mudflats and are considered to be a fern-ally.

Samples from Zone D exhibit a phytolith assemblage that suggests the continuation of coastal freshwater conditions; however, pond and surrounding marsh water levels may have been changed to levels more favorable to Tribe Oryzae grasses and freshwater sponge spicules. In fact, Zone D contains the highest relative abundance of freshwater siliceous sponge spicules. The spicules observed here were in good states of preservation and are not likely to have been redeposited through wind deposition or erosion. Freshwater sponges can occur in soils, streams and lakes, but may be more prolific in lake water depths of greater than 1 meter (Rader 1984), ponds and soils with higher organic matter content, and mesic soils (Schwandes and Collins 1994). This may indicate increased precipitation, a rising water table, or possibly less saline conditions for the pond and surrounding environments. Tribe Oryzae bilobate phytolith

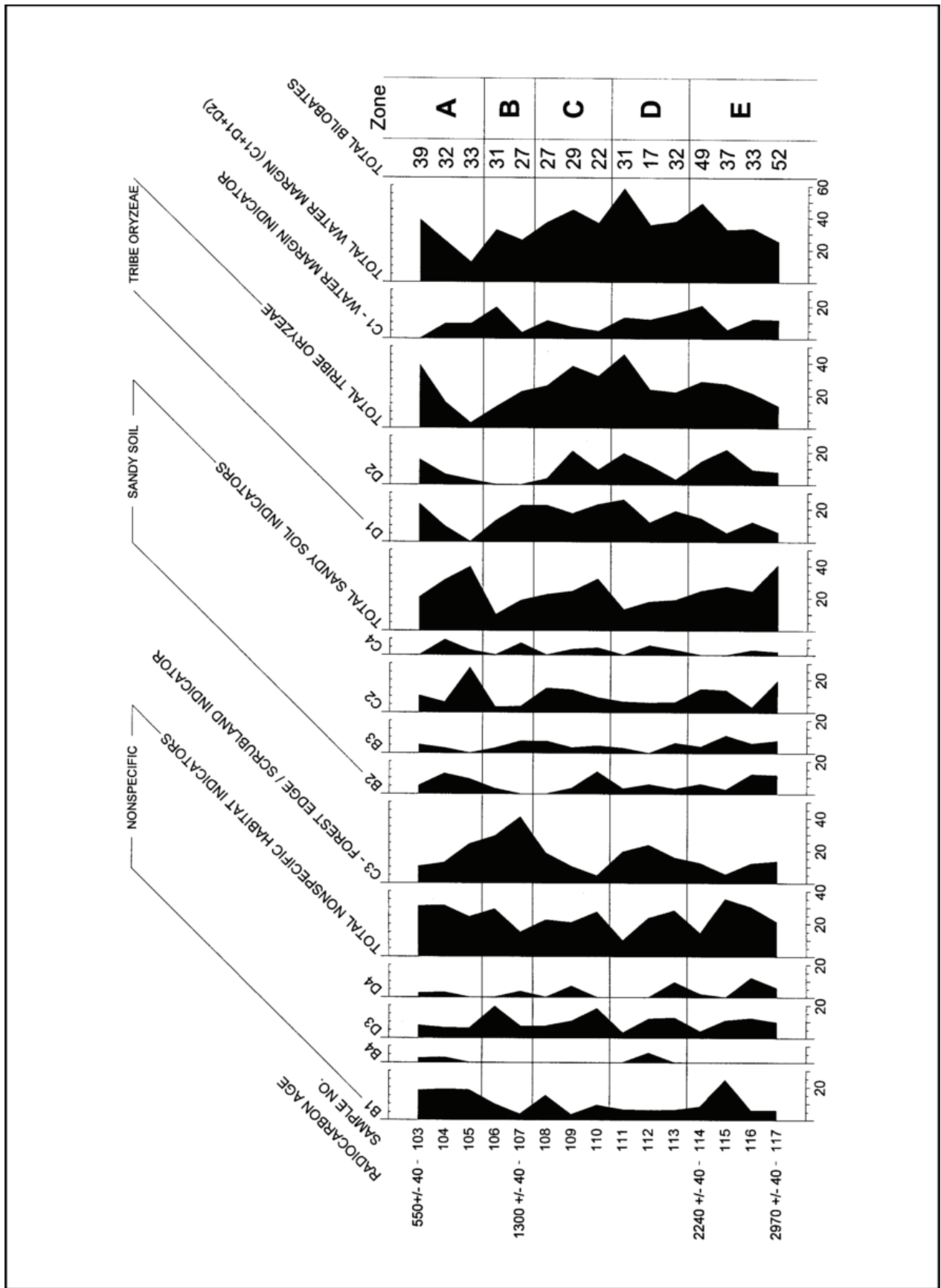


Figure A-3. Diagram of bilobate ecological indicators, Core 2, Site 41NU2, Cayo del Oso, Texas.

indicators reach their highest relative abundance at the top of this zone. *Leersia* prefers moist soils to shallow aquatic habitats. *Zizaniopsis* prefers shallow marshes, lakes and streams, and can tolerate somewhat brackish conditions. In general, tribe Oryzeae taxa can proliferate under conditions of periodic disturbance. Bilobate indicators of forest (scrub) edge and understory grasses are slightly elevated in abundance and may indicate increased precipitation. From the bottom to the top of Zone D there is an increasing abundance of fragmented faceted phytoliths, likely derived from tracheid elements of dicotyledonous trees and shrubs; however, they are indeterminate to any particular family. Some of these forms may also represent tracheid element fragments from the aquatic lycopod *Isoetes* (quillwort). *Isoetes* tends to be excluded from environments where temperatures exceed 20°C and precipitation is less than 60 cm per year (Sawada, et al. 2003). The pollen record (Figure A-1) contains its first occurrences of *Isoetes* in this zone, further supporting a slightly cooler and wetter environment at this time. Thus, the phytolith record for Zone D suggests that shallow freshwater pond conditions may have predominated, with periods of disturbance under slightly cooler, wetter, and perhaps stormier conditions than today.

Zone C predates the radiocarbon date of 1300 ± 40 RCYBP reported at the base of Zone B. This zone exhibits reduced quantities of Poaceae and *Quercus* pollen accompanied by increases in Cyperaceae, *Sagittaria*, and *Smilax*, representing sedges, arrowhead, and greenbriar. Although *Persicaria lapathifolium*-type pollen continues to be present, the frequencies are not quite as high as they were in Zone D. *Typha* pollen continues, as do *Isoetes* rhizoids. *Isoetes* microspores appear in sample 108 for the first and only time. The pollen record signals an expansion of the freshwater marshes.

Samples from Zone C exhibit a phytolith record dominated by a peak in occurrence of the indeterminate tracheid form discussed in the previous zone. Grass short-cell rondel phytoliths are at their lowest level of relative abundance, but this is, of course, a result of the high peak in tracheid elements and likely does not represent a decline in their absolute abundance. Wetland tribe Oryzeae phytoliths, both diagnostic morphotypes (bilobates, cross and rondels in Figure A-2) and those that are highly distinctive to diagnostic for the tribe (bilobates in Figure A-3), are still observed. Bilobate indicators of forest edge and scrubland grasses are at a very low level of abundance, concomitant with a fairly high occurrence of bilobate indicators of grasses that prefer sandy substrates. It is possible that this zone represents a climax in storm surge conditions. Lu and Liu (2005) demonstrated the feasibility of detecting the frequency of hurricane overwash deposition from inland lacustrine sediments. They were finding phytolith assemblages typical of dunes and outer bank islands within sand lenses from inland lake cores. It is possible

that the large sampling intervals (both within sample and between samples) employed here are smoothing (averaging) a record of increased storm frequency and intensity for Zones C and D. Deposition of the peak in indeterminate tracheid phytolith fragments may be related to plant accommodations to storm surge deposition. Radiocarbon dates supplied by the principle investigators for Core 2 bulk sediments were used to calculate approximate sedimentation rates (Table A-3) Zones C and D exhibited the highest rates of sedimentation (10.6 years per cm), thus adding further support for increased precipitation and/or storm surges. Thus, the phytolith record for Zone C suggests that shallow, slightly more brackish pond conditions may have predominated, with increased disturbance and storm intensity, under slightly cooler and wetter periods than today.

Zone B is anchored by a radiocarbon date of 1300 ± 40 RCYBP at its base. This zone notes a rise in Cyperaceae pollen and a concomitant decline in *Persicaria lapathifolium*-type, *Opuntia*, and *Sagittaria* pollen, suggesting drying conditions. *Typha* pollen increases and *Isoetes* rhizoids disappear from the record. *Smilax* pollen is present in sample 106, but not 107, indicating local growth of greenbriar at the forest margin.

Samples from Zone B exhibit a phytolith record suggestive of a slightly brackish marsh (Lu and Liu, 2005), as evidenced by an increase in smooth elongates, rondels and buliforms. Coastal freshwater marshes tend to have a greater representation of echinate elongates and bilobates, and salt marshes tend to have a greater representation of short-saddles from saltgrass (*Spartina alterniflora*). Tribe Oryzeae phytoliths are much reduced in abundance, as are bilobates associated with grass taxa that favor sandy soil substrates. Bilobates associated with grass taxa that favor forest edge and scrubland ecotones reach their highest levels for the entire record within this zone. The increase in smooth elongates and rondels actually starts within the topmost sample of Zone C. This rise is also analogous to the rise in *Persicaria lapathifolium*-type Polygonaceae pollen. Although *Persicaria amphibia*, the most hydrophytic of the native North American smartweeds, can occur in moist prairies to shallow ponds and is reported to be tolerant of slightly brackish conditions and as well as eutrophic conditions (Partridge 2001), it is not represented by *P. lapathifolium*-type pollen, as it has a distinct morphology type. This knotweed/smartweed is not known to grow in brackish conditions. The reduction in *P. lapathifolium*-type and *Sagittaria* pollen, might reflect more brackish conditions, since *Typha* pollen, representing cattail, which is capable of tolerating brackish conditions, increases slightly in this level. Thus, the phytolith record and other supporting evidence for Zone B suggests that slightly brackish marsh conditions existed at this time, possibly under slightly warmer and drier conditions than that for the previously discussed zones.

Table A-3. Data for Samples from the Cayo del Oso Site (41NU2)

Zone	Pollen Core 2	Phytolith Core 2	Sed Rate Core 2	Phytolith Core 4	Phytolith Core 6	Climate	Landscape
A	550 ± 40 Increasing oak, sedge marsh	550 ± 40 Drier brackish marsh	17.9 yr/cm	Sand dune w/ moist interdune areas		Increasing temp — ↑ Decreasing precipitation — ↓	↑ Increasing scrublands — ↑ Oak
B	Some oak, cattails, sedge marsh, shallow & stable	Brackish Marsh 1800 ± 40		1280 ± 40 Salt marsh			
C	Increasing sedges, shallow, dunes	Shallow brackish pond	10.6 yr/cm	Possible sand dune w/ moist interdune 1770 ± 40	Increased precip & Decreased temp 10.6 yr/cm (Sed rate)	More open	
D	Pond, more disturbance, fluctuating water depth, dunes	Freshwater pond		Sand dune w/moist interdune habitat		More open water, wetland	
E	2240 ± 40 Possible sedge marsh 2970 ± 40	2240 ± 40 Freshwater marsh 2970 ± 40	33.2 yr/cm	2670 ± 40 Possible freshwater marsh Dissolution	Fairly stable	More open, canopy thinning	
F	No samples	No samples		Fresh-brackish 3770 ± 40 swamp w/arboreal canopy	Wet & poss. cooler	Some canopy, increase in water table	
G	No samples	No samples		Maritime forest OR freshwater swamp w/ arboreal canopy	Warm	Closed canopy	

The pollen record for Zone A exhibits increases in *Quercus*, Chenopodiaceae, and Cyperaceae pollen, as well as reductions in *Persicaria lapathifolium*-type and Poaceae pollen. *Smilax* pollen rises slightly, then disappears at the top of the record, while *Typha* pollen remains relatively static in frequency. Sample 103 at the top of this zone is the only one to exhibit even a small quantity of *Prosopis* pollen, indicating that mesquite was present in the local vegetation. It is possible that mesquite joined this vegetation community relatively late. *Isoetes* appears in sample 105 at the base of this level. A radiocarbon date of 550 ± 40 RCYBP is reported for sample 103 at the top of the stratigraphic column. A very small quantity of microscopic charred Asteraceae tissue fragments are observed in sample 104, indicating burning herbaceous or shrubby plants at the site.

Samples from Zone A exhibit a phytolith record suggestive of a decreasing, slightly brackish marsh habitat, and an increasing area of drier habitats, possibly sand dunes, nearby. This is supported by a phytolith assemblage for Zone A which resembles that reported from modern Gulf Coast sand dunes samples (H. Lu and K.-b. Liu 2003; Lu and Liu 2005). Rondel phytoliths are at their highest levels of relative abundance here, and are common for both brackish marsh and sand dune habitats (Figure A-2). These rondels are likely derived from dry habitat grasses *Aristida desmantha*, *Uniola paniculata*, *Distichlis spicata*, *Sporobolus virginicus*, and tribe Oryzaceae marsh grasses *Leersia* spp. and *Zizaniopsis miliacea*. The slight increase in Chloridoideae short-saddle phytoliths are likely derived from the brackish marsh grass *Spartina patens* and *Eragrostis* spp. of sand dune areas. Bilobates associated with forest edge and scrubland grasses decrease rapidly throughout this zone, concomitant with some of the highest levels of bilobates associated with grasses that prefer sandy soil substrates (Figure A-3). Tribe Oryzaceae bilobates rise throughout this zone, possibly colonizing once-submerged areas of the marsh margins. A few globular echinate phytoliths diagnostic of the palm family (Arecaceae) were also noted in Zone A, and are likely derived from dwarf palmetto (*Sabal minor*) or Mexican palmetto (*Sabal mexicana*). Drier and warmer conditions for Zone A are also supported by the highest levels of charcoal and oak (*Quercus* spp.) abundance from the pollen record. Thus, the phytolith record for Zone A suggests increasingly dry and warm conditions at this time, resulting in a reduction of the slightly brackish marsh area and an expansion of dry habitats, most likely sand dunes.

Core 4 (dune deposits)

The phytolith record for Core 4 was severely affected by silica dissolution. Originally, a total of 23 phytolith samples were submitted for analysis. Ten samples underwent phytolith

extraction and were scanned for phytolith recovery (Samples 182 through 191). All of the samples were found to have experienced either mild dissolution (Sample 182) or severe dissolution (Sample 183 -191). Only Sample 182 yielded a reliable phytolith assemblage count (Figure A-4). Phytolith preservation and dissolution depends on 1) the particular type of phytolith and 2) the chemical and physical characteristics of the depositional environment (Piperno 2006). A third factor in preservation is the residence time in a particular environment. Relatively young calcareous sediments may still yield well preserved phytoliths. Thinly silicified phytoliths, particularly epidermal sheet elements and cell casts are readily dissolved in a wide variety of paleoecological contexts. More robust infillings of specialized plant silica cells, such as grass silica short cells, vascular tissue buliform cells and elongates are much more resistant to dissolution. In fact, buliform cells and elongates are often the only identifiable phytoliths in highly alkaline soils subjected to periodic moisture. In extreme cases, such as perennially moist calcareous soils, phytoliths will often be completely dissolved, with no recovery possible.

The samples from Core 4 varied in color from light gray to grayish black, all of which with evidence for varying degrees of calcium carbonate nodule formation. All of the samples were moderately to severely reactive when tested with hydrochloric acid, indicating the highly calcareous nature of these dune deposits. The phytolith record from Sample 182 exhibits an assemblage with characteristics fairly analogous to that reported by Lu (2005) from some of their modern Gulf Coast sand dune habitats. After two hours of scanning, only five phytoliths were recovered from Sample 183. Evidence for severe dissolution was observed in the form of large dissolved pits completely covering the phytoliths observed. As a result of this evidence and the high reactivity of the remaining samples in HCL, Samples 184 through 191 were scanned to assess the likelihood of phytolith recovery. It was determined that the remaining samples would not yield enough phytoliths to provide an unbiased and statistically significant assemblage. Although Sample 182 was highly reactive in HCL, residence time was apparently not long enough to severely affect phytolith preservation. Sample 182 was recovered from a depth 40 to 43 cm. The closest date comes from the 65 to 68 cm level, and is reported as $1,060 \pm 40$ years. An estimate for the age of Sample 182 would be approximately 500 RCYBP. This would place the sample either above or within Zone A previously discussed. Sample 182 likely represents a sand dune or wet meadow interdune environment with a relatively brackish and shallow watertable. The vegetation was likely dominated by dune grasses *Uniola paniculata* and *Aristida desmantha*. Despite poor phytolith preservation, it is likely that pollen is preserved in these sediments and would be recoverable by Paleo Research Institute.

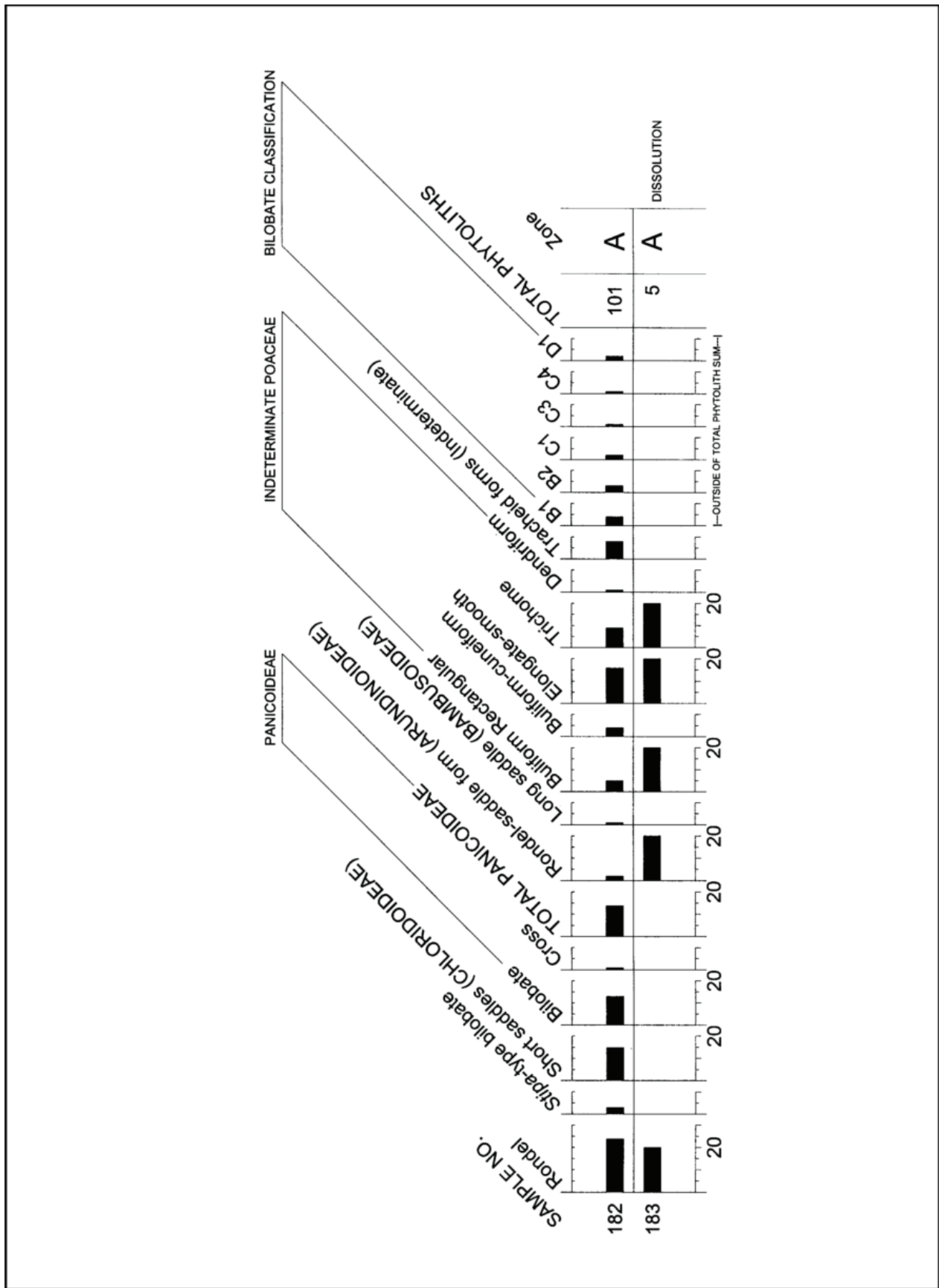


Figure A-4. Phytolith diagram for Core 4, Site 41NU2, Cayo del Oso, Texas.

Core 6 (dune deposits)

The phytolith preservation and recovery from Core 6 sediment samples was mixed. A total of 16 samples were submitted for analysis, 4 of which yielded statistically insignificant phytolith counts, two of which yielded no phytolith recovery. Because of the mixed recovery record and the disjunct nature of statistically significant counts, a bar histogram rather than a silhouette-style graph was used (Figure A-5). Likewise a separate bilobate classification graph was not created. Most of the sediments contained calcium carbonate nodules. Samples exhibiting low reactivity with HCL yielded good phytolith preservation (279, 280, 287). Samples exhibiting moderate to high reactivity with HCL yielded little to no phytolith recovery. The one exception was Sample 294, which was highly reactive to HCL, but exhibited good phytolith preservation with some evidence of dissolution. It is possible that Sample 294 phytoliths preserved as a result of limited exposure to moisture. Zonation for Core 6 was primarily based on chronological and assemblage association with that for Core 4. Older and non-analogous assemblages were given additional zone designations.

Sample 294, the only sample from Zone G, exhibits a phytolith assemblage that does not fit well with any of the Gulf Coast habitat phytolith assemblages described by Lu and Liu (2003, 2005). The closest analogs are described from *Juncus*/sedge fen, maritime forest, and freshwater lake habitats (H. Lu and K.-b. Liu 2003; Lu and Liu 2005). Globular echinate phytoliths diagnostic to the palm family (Arecaceae) were the most abundant form, and are likely derived from dwarf palmetto (*Sabal minor*) and/or Rio Grande palmetto (*Sabal mexicana*). Dwarf palmettos grow to 4 or 5 feet, require part sun to shade, and prefer neutral to alkaline soils of swamps and river bottoms. Rio Grande palmetto, also called Texas palm, grows to 50 feet, requires full sun to part shade, and can adapt to moist or dry soils. The presence of either species suggests that forest-type vegetation was dominant in the area at this time. Indeterminate grass phytolith buliforms and elongates are also very abundant. Although phytolith preservation was generally very good, evidence for dissolution was observed. It is likely that lightly silicified faceted phytoliths diagnostic of leafy dicot trees and some grass silica short cell phytoliths were lost to dissolution. Maritime forests and swamps are comprised of a number of tree species and understory grasses from the Pooideae and Bambusoideae subfamilies. Trapeziform sinuates and long saddles, diagnostic of the Pooideae and Bambusoideae respectively, were observed, but in low abundance, possibly due to dissolution. Although no date is closely associated with this sample, it is likely older than 4,000 RCYBP, and at a time when sea levels may have been a meter or more lower. Thus, the phytolith record for Zone G suggests that a

maritime forest, or freshwater swamp with an arboreal canopy was the dominant vegetation type at this time.

Samples from Zone F exhibited evidence of severe silica dissolution; however, some phytoliths were well preserved, or at least identifiable. As a result, preservation may have been biased towards certain morphotypes, and may account for the high abundance of indeterminate tracheid fragments. As discussed for Core 2 samples, the possibly fractured nature of this morphotype makes differentiation between monocot (grass) and dicot (tree, shrub, forb) xylem ambiguous. Despite this, enough phytoliths were observed to provide an informative and interpretive assemblage. Zone F exhibits characteristics similar to Zone G, but overall vegetation species richness may be greater. Trapeziform sinuates, long saddles, and globular echinates, diagnostic of the Pooideae, Bambusoideae and Arecaceae respectively, were observed and suggest that some degree of a forested canopy with a palm component existed. However, the presence of tribe Oryzeae and Cyperaceae (*Juncus*/sedge) phytoliths, as well as diatoms and sponge spicules, suggest that moisture levels may have been higher than that experienced for Zone G. This may be the result of a shallow watertable, or the presence of open water. The upper portion of Zone F has an associated radiocarbon date of 3770 +/- 40 B.P. Thus, the phytolith record for Zone F suggests that a fresh/brackish swamp with an arboreal canopy was the dominant vegetation type at this time.

Samples from Zone E exhibited evidence of severe silica dissolution, with little to no phytolith recovery. The upper portion of the zone has an associated radiocarbon date of 2760 +/- 40 B.P. The phytolith assemblage from Sample 289, and the chronological association of the zone suggests an expansive freshwater marsh habitat affinity with Zone E from Core 2 (Figure A-2); however, the likely biased nature of preservation/dissolution makes an accurate interpretation uncertain.

Sample 287 represents a single sample that may share a chronological affinity with Zone D from Core 2, interpreted as a freshwater pond habitat. However, a lack of diatoms and sponge spicules, and a higher percentage of sandy soil bilobate indicators (not graphed separately) suggests that this location (Core 6) may have been drier than Zone D from Core 2. The presence of a few tribe Oryzeae phytoliths suggests that these may have been moist soils. A few globular echinates diagnostic to the palm family (Arecaceae) were also noted and could possibly be derived from Texas palm (*Sabal mexicana*). The upper boundary of this zone has an associated radiocarbon date of 1,770 +/- B.P. Thus, the phytolith record for Zone D suggests the presence of a sand dune habitat with a shallow watertable at this time.

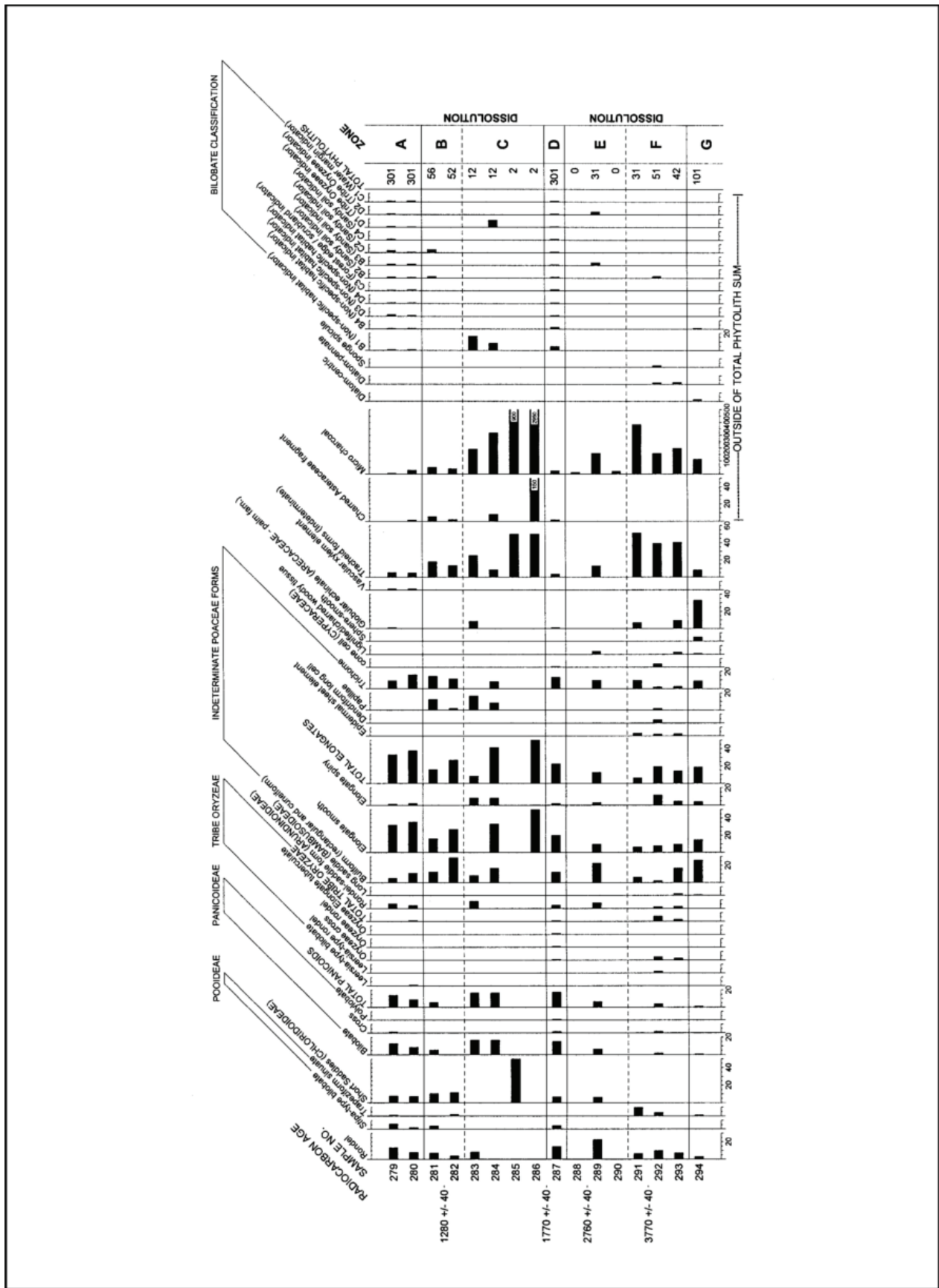


Figure A-5. Phytolith diagram for Core 6, Site 41NU2, Cayo del Oso, Texas.

Samples from Zone C exhibited evidence of severe silica dissolution, resulting in very little phytolith recovery. Samples from this zone share a chronological affiliation with Zone C from Core 2, a shallow brackish pond or marsh; however, the upper boundary designation is somewhat uncertain (Figure A-5 dashed line). Sample 283 from the top of the zone yielded a few phytoliths diagnostic of the Arundinoideae, and likely derived from common reed (*Phragmites australis*), an obligate wetland plant tolerant of slightly brackish conditions. The poor phytolith recovery from Zone C sediments make an accurate interpretation very difficult; however, the presence of common reed, and the likely existence of moist sediment conditions necessary for phytolith dissolution may indicate the existence of moist sand dune/interdune conditions at this time.

Samples from Zone B exhibited evidence of moderate silica dissolution, but not enough to make phytolith identification impossible. This zone shares a chronological affiliation with Zone B from Core 2, interpreted as a slightly brackish marsh. A radiocarbon date of 1,280 +/-40 B.P. is associated with this zone. The phytolith assemblage, as evidenced by the high proportion of saddles and buliforms and the low proportion of rondels and bilobates (Lu and Liu 2005), and the absence of common reed (*Phragmites australis*) suggests that a *Spartina alterniflora* dominated salt marsh existed here at this time.

Samples from Zone A yielded phytoliths that were well preserved, with no evidence of dissolution. This zone shares a chronological affiliation with Zone A from Core 2, interpreted as a contracting brackish marsh habitat experiencing increasing temperatures and decreasing precipitation. The phytolith assemblage here shares characteristics for both dune and interdune meadow habitats as reported by Lu (2003; 2005). The phytolith assemblage here suggests a dune habitat with an elevated watertable, as evidenced by a high proportion of bilobates and rondels, and a moderate amount of short saddles. The bilobate classification assemblage (not graphed independently) is dominated by sandy soil indicators. The presence of phytoliths diagnostic of *Phragmites australis* (Arundinoideae) and palm (Arecaceae), suggest the presence of moist, interdune meadows nearby. It appears that this area may have been experiencing a drying trend, starting from the upper level of Zone B (Sample 281), and proceeding through Zone A, as salt marsh converted to a drier sand dune environment. This warming and drying trend seems to be analogous to that reported for Zones B and A from Core 2.

Summary and Conclusions

Companion pollen and phytolith analysis of Core 2 yielded evidence of local and regional vegetation that included a variety of trees such as pine, juniper, alder, birch, a member

of the sycamore family, mesquite, oak, and elm. The most visible elements of the local vegetation included grasses, various members of the sunflower family, members of the Chenopodium group, and sedges. Beyond this, smartweed/knotweed, prickly pear cactus, and cattails also would have been visible. Mustards might well have colored the vegetation community during the spring when they flower. Arrowhead was important locally intermittently, primarily between approximately 2240 and 1300 RCYBP. Poison ivy is present in the record, but not well represented, which is often the case, as this pollen tends to be under represented. Dominant grasses varied as dune/interdune and marsh areas expanded and contracted in response changing moisture availability and levels of salinity. Beach, dune and interdune areas were likely dominated by grasses *Panicum amarum*, *Uniola paniculata*, *Aristida desmantha*, *Cenchrus* spp., *Spartina* spp., and occasional clusters of dwarf palmetto and possibly Texas palm. Pond and marsh habitats were likely dominated by *Leersia* spp. and *Zizaniopsis miliacea*, with *Phragmites australis* and *Spartina* spp. also present. Forest understory and shaded areas of forest edge and scrubland were likely occupied by *Panicum* and subfamily Pooideae and Bambusoideae species.

The phytolith record from Core 4 was severely affected by silica dissolution, likely due to perennially moist conditions of calcareous (alkaline) soils. These conditions are well known to dissolve many, and sometimes all phytolith types. At a minimum they tend to present a biased record of phytolith occurrence. The area represented by Core 4 was likely a sand dune and moist interdune environment with a shallow brackish watertable. The vegetation would have been dominated by grasses and other dune plants. Despite poor phytolith preservation, it is likely that pollen is preserved in these sediments and would be recoverable by Paleo Research Institute, thus providing a better record of the sand dune vegetation community.

Although the Core 6 phytolith record was also affected by silica dissolution, enough phytoliths were recovered to provide an accurate interpretation. The area represented by Core 6 sediments provided some of the most dramatic evidence of long-term habitat change. Initially a landscape dominated by a closed canopy forest of trees, shrubs and vines (> 4,000 RCYBP) gave way to slightly more open marsh-type vegetation of shrubs, grasses, sedges and dwarf palms (≈2670 RCYBP), and then opened-up to a drier and salt tolerant grass and herb dominated sand dune and interdune environment (≈800 RCYBP).

A tabular summary of habitat, climate and landscape-level change spanning the last approximately 4,000 RCYBP for the Cayo Del Oso site is presented in Table A-3 and summarized

here. An expansive maritime forest or freshwater swamp dominated by a closed arboreal (possibly palm dominated) canopy at approximately 4,000 RCYBP opened-up slightly to a fresh or brackish swamp with dwarf palms, shade tolerant grasses and aquatic emergent grasses in more open areas. By around 3,000 RCYBP, the shady swamp habitat gave way to a more open marsh habitat with emergent aquatic and water margin grasses proliferating. Arboreal trees such as oak and elm were still present, but perhaps more scattered. This marsh may have been expansive, possibly due to a slight dip in sea-level. Dune areas may have also been expansive at this time of possibly warmer and drier conditions. At approximately 2,200 RCYBP, precipitation, sedimentation rates, and possibly storm frequency/intensity increased. This increase in precipitation may have resulted in freshwater pond development from the previous marsh habitat. A rising water table may have caused less tolerant trees to give way to wetland plants and other plants adapted to colonizing disturbed areas along marsh and pond margins such as *Isoetes*, *Persicaria*, *Sagittaria*, *Phragmites*, *Leersia*, *Zizaniopsis* and members of the Apiaceae and Cyperaceae families. Sandy soil areas closer to the bay and ocean may have lost some stabilizing vegetation, resulting in increased dune activity and annual plants such as members of the Chenopodiaceae and Brassicaceae families more adapted to frequent disturbance events. This environment may have persisted until around 1,300 RCYBP, when the apparent disturbance period subsided. Precipitation may have decreased and been accompanied by a slight increase in temperature. A lowering of the water table may have caused pond habitats to convert to shallow brackish water marsh habitats. Cyperaceae taxa such as sedges and rushes became increasingly abundant, as do cattails. The cattail pollen is likely derived from *Typha domingensis*,

a wetland plant that prefers brackish or nutrient-enriched wetlands. *Typha* is also not very tolerant of rapidly fluctuating water levels and further indicates a more stable environment after 1,300 RCYBP. Generally, this habitat persists for the remainder of the pollen and phytolith record which ends at approximately 550 RCYBP; however, tribe Oryzaceae grasses (*Leersia* and *Zizaniopsis*), oak trees (*Quercus*), pine trees (*Pinus*) and mesquite (*Prosopis*) increase towards the top of the record which ends at approximately 550 RCYBP. Oaks, pines and mesquite are perhaps becoming more competitive from increased temperature, periodic wild fires and drier conditions. Oryzaceae genera are perhaps colonizing once saturated marsh habitats.

Paleo pond sediments from the Cayo Del Oso site present a unique opportunity for phytolith and pollen paleoenvironmental reconstruction for coastal South Texas. The phytolith and pollen record indicates that there has been some subtle, long-term environmental change as well as a few punctuated events and/or short term changes strong enough to be detectable with the sampling interval strategy employed. The large sampling intervals, both within and between samples, limits the ability to detect punctuated events and has an overall smoothing effect on the environmental record. In particular, additional phytolith sampling of smaller interval pond sediments has the potential to detect hurricane storm frequency for South Texas. Also, there is a high likelihood that pollen is preserved in dune and interdune deposits and recoverable by personnel at Paleo Research Institute. Analyzing additional samples from these deposits could greatly enhance temporal resolution of the paleoenvironmental record for Cayo Del Oso and coastal South Texas.

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APPENDIX B:

Pollen Analysis Of Pond Core Samples: Oso Bay Site, Nueces County, Texas

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Introduction

Data Base

The current study focuses on 24 pollen samples collected from a sediments core extracted north of the Cayo del Oso Site (41NU2). The archaeological site is located on the shores of the False Oso Bay and is only about 50 meters from Corpus Christi Bay. Various cores were collected from a pond north of the site using an Eijkelkamp percussion drilling set operated by Charles Frederick. The cores were later split open in the lab, samples were extracted, and the dirt from 23 individual levels was sent to us for processing and fossil pollen studies (Table B-1). No surface soil control samples were collected.

Pollen Extraction

There are a number of articles, chapters, and even books detailing the methods that one can use to successfully extract fossil pollen from various types of sediments (Hunt 1985; Riding and Kyffin-Hughes 2006), including some articles that focus specifically on techniques used for archaeological sediments (Bryant 1988; Bryant and Holloway 1983; Coil et al. 2003).

Pollen extraction is a critical part of any project because the use of the wrong acids and methods can lead to the loss of fossil pollen, while other methods may not remove enough of the detritus to permit accurate identification of the fossil pollen in the remaining residue. For these reasons it is essential that reports, such as this one, include the precise methods used during the extraction procedure. In this way the reader may be assured that pollen recovery was maximized and that fossil pollen was not inadvertently destroyed or lost during processing.

During the work for this project, all procedures were conducted under a fume hood using sterile, surgical gloves in our sealed Palynology Laboratory. In addition, glycerine-coated slides are routinely left exposed in various locations within the lab. These slides are checked weekly for any signs of outside pollen contamination. None was noted during or after this project was being conducted. Thus, I am certain that none of the pollen I found in these samples came from contamination in

our facility. The extraction procedure I used for these samples is consistent with procedures detailed in many reports and ones that we have used successfully for several decades. The precise processing steps and the sequence in which they were used on all samples are listed below:

1. From each of the core samples I removed 20 grams of soil. The soil from each sample was placed in a sterile 800 ml plastic beaker. Next, I added 50 ml of concentrated HCl to dissolve any calcium carbonates in the samples. At that time I also added two Lycopodium tablets to each of the samples. Each Lycopodium tablet contained 13,500 tracer spores. The use and amount of tracer spores that should be added to archaeological and other types of pollen samples has been extensively studied and reviewed by Louis Maher (1981). He notes that the ratio of tracer spores added to samples should be in a ratio of between 1:1 and 1:2 (tracer spores vs. fossil pollen) in order to achieve statistically accurate calculations of fossil pollen concentrations per gram or milliliter of sediment. The ideal, according to Dr. Maher, is to try to achieve as close to a 1:2 ratio as possible without exceeding it. Once the ratio of tracer spores vs. fossil pollen exceeds the 1:2 ratio significant statistical errors often occur, as discussed by Jones and Bryant (2007). I use Lycopodium spores as a tracer in most sediment studies from sites located in the temperate and arid regions of the American Southwest and Texas because it is extremely rare to find naturally occurring Lycopodium plants growing anywhere in those regions. Therefore, the only potential sources of these spores come from those that have been added.

2. After all reaction with the HCl had stopped, I filled the beakers with distilled water and allowed the beakers to stand for a period of four hours. After that, the liquid portion was carefully siphoned off. Pollen will remain suspended in water for various periods depending on the temperature and salinity. Nevertheless, after four hours, all fossil pollen will settle to the bottom of a container and thus the liquid portion can be carefully poured or siphoned off without loss of fossil pollen, provided the sediment at the bottom of the beaker remains undisturbed (Lentfer et al. 2003). The remaining sediment in each beaker was then poured into 50 ml centrifuge tubes (CT) and spun down for 3 minutes at a speed of 2,500 rpm. The liquid was then

Table B-1. Pollen samples collected and processed from core sediments extracted from the pond located north of the Cayo del Oso Site (41NU2)

Core Sample Numbers processed:

**88, 89, 90, 91, 92, 93, 94, 95, 96, 97, 98, 99,
100, 101, 102, 178, 179, 180, 181, 275, 276,
277, and 278**

poured off. This process was repeated until all sediment from each beaker had been concentrated in each CT.

3. The collected sediment from each centrifuge tube was then placed into a clean 800 ml plastic beaker. Because these sediments contained a high clay content, I needed to disperse and deflocculate the sediments. This step is essential because fine-grained minerals, such as clays, may form clumps that trap pollen and spores inside. Often, clumps form because particles in suspension have atoms of unsatisfied valences, thus they attract other ions and form clumps. There are a number of dispersing agents that tend to work effectively as a deflocculant. I have found that one of the best is a 5% solution of Calgon water softener containing sodium pyrophosphate ($\text{Na}_4\text{O}_7\text{P}_2$). Not all types of Calgon contain sodium pyrophosphate so one must choose the correct one. Once the 5% solution was mixed, I added 700 ml to each beaker, stirred them once each hour for four hours and then let the samples sit over night without being stirred again. The next day after all the clay had been dispersed, the liquid portion was siphoned off without loss of fossil pollen.

4. Next, I added 100 ml of concentrated HF (58%) to the sediment in each beaker, stirred the sample, covered them to prevent potential contamination and let them sit overnight in the fume hood. I might caution others that this is very dangerous and that the HF must be added very slowly at the rate of only 1-2 ml at a time. The sample must then be thoroughly stirred and allowed to sit for one minute before adding more HF. After HF is added, it will slowly heat the sediment as it begins to react with the silicates. Some types of sediment, especially those containing fine-grained clays and mica, will react violently to the HF. This reaction, however, does not occur at first, it occurs only after the sediment reaches a high temperature above 150oF. If one adds too much HF at first, once it heats, the HF reaction occurs so rapidly that the liquid will bubble rapidly and can flow out of the beaker. I use HF only with the greatest degree of caution because it can cause permanent injury or in extreme cases can cause death to an individual. Medical reports state that if 2% or more of one's body comes in direct contact with liquid HF, the result is always fatal. There is no adequate antidote for that type of HF exposure. The HF process removes most of the fine-grained silicates from the sample and does not damage the pollen. After stirring the HF in each sample they were covered and allowed to sit in the fume hood overnight. The next day I filled each beaker with distilled water, stirred each sample and let it sit for four hours before siphoning off the liquid. Once the remaining liquid was spun down in a 50 ml CT, each sample was rinsed twice with distilled water, and then the CT was filled with concentrated HCl. This HCl step is necessary to ensure removal of potential fluorosilicates in

each sample, which often form during the HF treatment. Sometimes two or three HCl rinses are needed before all fluorosilicates are dispersed and the samples can be rinsed again in distilled water and neutralized.

5. I then transferred the material in each sample to a 15 ml CT and added 10 ml of 5% KOH. Each sample was stirred and then heated in a heating block for 10 minutes at 180oF. This was followed by two washes in distilled water. Next, I added 10 ml of concentrated HCl and heated each sample for one minute. I then spun down the sample, poured out the liquid, and again rinsed the sample twice in distilled water. This HCl step is essential to remove any remaining humic acids and dissolved compounds that might be present in each sample after completing the KOH treatment.

6. Each sample was then rinsed in glacial acetic acid, centrifuged, and the glacial acetic acid was carefully poured off.

7. Next, I added 10 ml of a mixture of 1 part sulfuric acid to 9 parts acetic anhydride. This is known as the acetolysis process (Erdtman 1960). After heating each sample in a heating block at 180oF for 10 minutes, I centrifuged the solution and then poured off the liquid.

8. I then rinsed each sample in glacial acetic acid and after centrifuging, I poured off the glacial acetic acid. I want to warn others that it is essential to rinse the sample in glacial acetic acid both before and after the acetolysis process. The acetolysis mixture does not mix with water and will react violently if it comes in contact with any water. The glacial acetic acid will remove excess water from samples.

9. Next, the samples were rinsed twice in distilled water.

10. The next step was to fill each CT one-half full of zinc bromide having a specific gravity of 2.0. That solution was thoroughly mixed with each sample for 30 seconds on a vortex stirrer to ensure complete mixing of all solid material in the CT. The samples were then placed in a test tube rack and allowed to sit for 5 minutes. After that, I then spun the samples at 500 RPMs for 5 minutes. This was followed by a spin of 5 minutes at 2,500 RPMs. I find that by doing the separation using these three steps (one resting and two stages at two different centrifuge speeds) I get a much better separation of pollen from the mineral and remaining silicate detritus in samples. When this was completed, I very carefully pipetted off the upper layer of the zinc bromide, which contained all of the pollen. The zinc bromide solution containing the pollen was placed into a 100 ml beaker and 80 ml of 95% ETOH was added to reduce the specific gravity and permit the pollen to sink during centrifuging. All of the solution from each sample was carefully spun down in CTs.

11. I then repeated step 10 using the sediment that has not floated during step 10. This second zinc bromide treatment ensures that any remaining trapped pollen and spores will be liberated and allowed to float to the surface and be removed at the end of the procedure. I then checked the material that remained at the bottom of each CT to ensure that no pollen or tracer spores were lost after completion of this second procedure. None was found.

The final step consisted of rinsing the residue in each sample twice in water, twice in ETOH, adding two drops of safranin-0 stain, stirring the sample, and finally rinsing each CT once more in ETOH. The remaining residue was then pipetted into a one-dram, glass vial with a screw top. Five drops of glycerine was added to each sample and the samples were placed on a warming plate to enable the remaining ETOH to evaporate overnight.

Pollen Counting and Identification

Pollen Counting

When the remaining ETOH had evaporated in each of the processed samples, I carefully stirred the residue in each vial and added additional glycerine, if needed, to thin the solution. I then prepared several individual slides from each sample using the technique explained in the paper by Jones and Bryant (2001). As they explain and demonstrate through testing, incorrect methods of sub sampling can skew pollen counting results from individual samples. Pollen counting was performed using a Nikon Optiphot binocular microscope at magnifications ranging from 400x-1000x and all photographs were taken at 400x using a Nikon 950 Coolpix camera attached to a trinocular head. Identifications of pollen and spore types in each sample were checked against modern pollen reference materials, which now number more than 20,000, and are on file in the Texas A&M Palynology Laboratory. These modern pollen reference materials include the Texas A&M Modern Pollen Reference Collection, the Mobil Oil Modern Pollen Reference Collection, the Meredith Lieux Modern Pollen Reference Collection, and the AMOCO Modern Pollen Reference Collection.

How Many Pollen Grains to Count

How many pollen grains one should count in a sample has been the subject of much debate among palynologists (Barkley 1934). For pollen wash samples one usually counts all of the recovered pollen, unless there is an excessive amount of pollen and residue recovered. For all types of non pollen wash samples I usually count between 200-300 pollen grains because counts of this size have been shown to be adequate for these types of studies. As noted in Traverse (2007), the standard deviation in

a 100 grain pollen count for a pollen type that might appear in a normal percentage as low as 1% is 100%, thus the probability of finding that pollen type in any given 100 grain pollen count from the same sample would be between 0-2% in two-thirds of the additional pollen counts from that one sample. Even if pollen counts were increased to 1,000 pollen grains per sample, a pollen type occurring in a normal percentage of 1% would still have a standard deviation of 30%, meaning that in two-thirds of all additional 1,000 grain pollen counts of the same sample that pollen type would occur within the range of .3-1.3%. In terms of time and costs, the higher pollen counts are not normally justified, unless they are considered critical and worth the added cost. Jones and Bryant (2001) conducted experiments to test this hypothesis and discovered that the only significant thing usually gained by increasing to larger pollen counts above 200 grains per sample is the potential for finding additional, rare taxa, which normally occur in very low frequencies. Furthermore, as they showed in their experiments, even with high pollen counts, the percentages of the primary pollen taxa in a sample remain statistically similar in repeated counts at the 200 grain level, provided each count came from slides that were sub sampled correctly (Jones and Bryant 1998).

Pollen Identification

Pollen identification of fossil material in these samples was often quite difficult to determine because of the high degree of pollen alteration and degradation caused by unfavorable depositional conditions at the site where the cores were extracted. As noted in the pollen micrographs (Figs. 1-5) of fossil material examined from these samples, many of the pollen grains are so altered that they only vaguely resemble their original shapes or ornamentation. The highly degraded pollen grains that could not be identified with any degree of certainty are listed in the pollen counts as "degraded." They are included in the pollen counts of all samples because accurate pollen concentration values can only be derived when "all" of the fossil pollen is recorded whether or not it can be identified to family, genus, or species. Cushing (1966) was the first to test and report that a failure to include all degraded pollen in the actual pollen counts will produce unreliable and inaccurate pollen concentration values. To calculate fossil pollen concentration values in these and all other samples one must know three points of reference: 1) how many tracer spores have been added to the sample and how many were counted during the actual pollen count, 2) exactly how much sample material was used in each sample and how was it measured in grams or cubic centimeters, and 3) how many fossil pollen grains were counted in each sample including both identifiable and degraded pollen.

Fossil pollen preservation in all of these samples was marginal at best. Because I have been examining these types of highly

degraded fossil samples from sites throughout Texas, the Middle East, and the American Southwest for more than 40 years I have learned to recognize and correctly identify many highly degraded examples of pollen, which might not be recognized by others fairly new at working with these types of deposits. Regardless of how badly degraded some of the fossil pollen appeared in these samples, if I could be certain of the grain's original identity, I counted it as a pollen grain in that taxon, not as a degraded pollen grain. Only the pollen grains in these samples that were beyond any possibility of

identification were listed as degraded. Some examples of those are illustrated in Figure B-1.

Because the pollen in these samples were so badly degraded, if I found a fragment of a pollen grain that represented more than approximately one-fourth of the original pollen grain, I counted that fragment as a whole grain of that taxon. Individual pine bladders, or the body of a pine pollen grain were each recorded as one-half a pine pollen grain.

Scale for all photos: Space between numbers is equal to 25 microns.

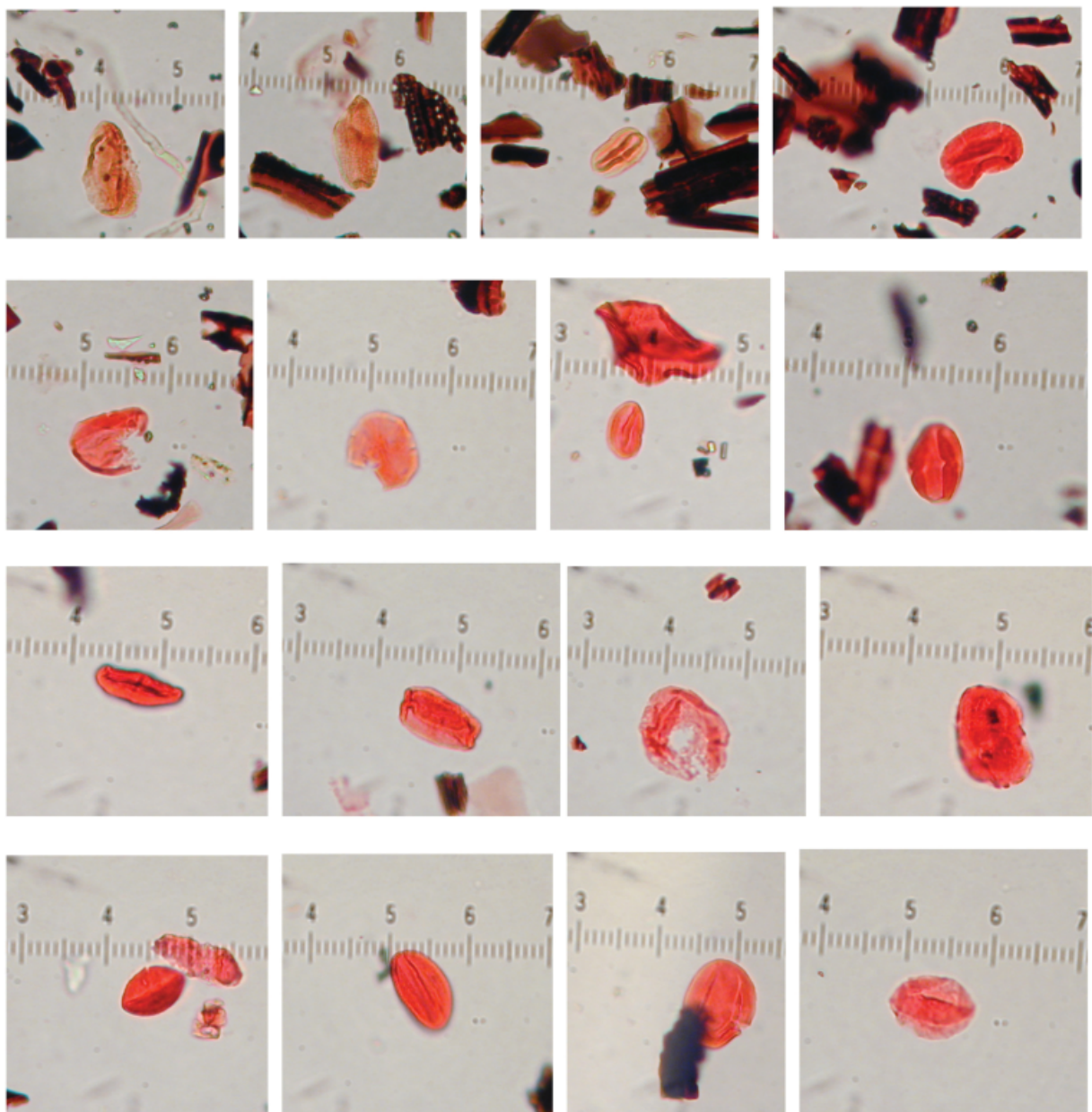


Figure B-1. *Examples of Degraded Pollen Beyond Recognition in the Core Samples.*

During the pollen counting of these samples I had to make important decisions in terms of how to combine certain categories of fossil pollen. For example, the Asteraceae plant family contains more than 1,500 genera and more than 22,000 species that grow in almost every known habit (Mabberley 1997). Asteraceae pollen can be divided into certain categories, most of which occurred in these samples. One primary group, which is insect-pollinated, is called the “high-spine” group because their pollen grains have a surface morphology consisting of long spines greater than 2.5 microns in length (Martin, 1963). These pollen types are very durable and even when degraded they can often be recognized (Fig. 2). Three other major pollen groups within the composite family are fairly easy to separate into distinct categories. The pollen types that are similar to and include ragweed are all wind-pollinated and are assigned to the category called “low-spine” types because their surface spines are shorter than 2.5 um in length (Martin, 1963). Another group of composites have fenestrate pollen grains and belong in the tribe Lactuceae (Fig. 5). The final group containing *Artemisia* is a group of wind-pollinated composites that all have a very distinct type of morphology and thus can be separated into this one genus. There are other pollen types produced by plant genera within the composites that are also distinctive enough that they can often be identified and listed separately as a genus, provided they are not severely degraded. Several of these other taxa include *Centaurea* (star thistle), *Cirsium* (thistle), and *Mutisia* (mutisia). For most of the genera and species in the Asteraceae family their pollen morphology is not distinctive enough to warrant separation into a genus or species without extensive regional pollen keys produced at the resolution level of a scanning electron (SEM) or transmission electron microscope (TEM).

Paul S. Martin (1963), while examining deposits from sites in the American Southwest was the first to propose the term “Cheno-Ams” for the combined groups of pollen in the family Chenopodiaceae and the genus *Amaranthus* in the Amaranthaceae family. Pollen grains in both of these groups are nearly identical in appearance because all the types are similar in size and shape and all are periporate (having many pores on the surface). These characteristics generally make the genera in this group indistinguishable using LM (light microscopy), unless extensive regional studies of specific taxa in these plant groups have been conducted and the pollen grains categorized according to their cord/diameter ratios (McAndrews and Swanson 1967). I have followed the use of Martin’s category of Cheno-Ams for these types of pollen grains found in the samples of this study.

The genus *Pinus* (pines) contains more than 90 species that can be divided into two major groups based on various

characteristics including their leaf morphology and their pollen types. In pine pollen the presence or absence of verrucae on the underside (distal surface) of the pine pollen grain in the germinal furrow region between the bladders will identify to which of the two major categories the taxon belongs. The two major groups of pines were first separated into “diploxylon” (those species with pollen having no distal verrucae) and “haploxylon” (those species having distal verrucae) pollen categories by Ueno (1958). In this present study I found the majority of the pine pollen was of the diploxylon variety. I did not find any grains that were of the haploxylon type. In a few cases where the area of the germinal furrow was obscured, I assumed those grains were also of the diploxylon variety and thus added them to that category. Likewise, when individual pine bladders were found, they were also counted as coming from diploxylon pines. If one finds that my category assumption for all the pine pollen in these samples is troublesome, then one should consider all of the pine pollen as being unclassified as to specific pine category.

Ephedraceae (joint fir) pollen occurred as single grains in only two of the samples in this study (Sample 95 and 101). This plant family contains only one genus with about 65 different species. The pollen in the single genus *Ephedra* (joint fir) can be divided into four major groups (Steeves and Barghoorn 1959). Two of the four major groups are common in North America and the pollen grains in both groups look similar, yet each has some unique morphological characteristics. The first of these two, the *Ephedra nevadensis* group, contains a number of individual species. Each of those species produces pollen with fewer furrows and many of the raised ridges, called plica, have furrows with hyalines. The second of these two is called the *Ephedra aspera* group and it also contains a number of individual species. The species in this second group produce pollen grains that have more plica and furrows but none have hyalines. Although both the joint fir pollen grains found in this study were degraded, both appear to come from the *E. aspera* group (Fig. 5).

Pollen Rain and Pollen Preservation

The pollen rain includes all of the pollen that is dispersed and deposited at a given locale. One can examine and discuss the pollen rain of a large geographical region, or focus on the pollen rain deposited in a very tiny geographical area, depending on the objectives of the pollen study. At any given location, such as location of the cores in this study, the eventual interpretation of recovered pollen data must account for all factors that may have influenced the composition of the original pollen rain, and later for the factors that may have affected and altered the initial composition of the buried pollen assemblage.

During the last 50 years, palynologists have learned there are many complex factors that determine the original pollen rain in a region. These include factors such as: 1) types of pollination; 2) differences in pollen production; 3) differential dispersion patterns; and 4) the size, weight, and aerodynamic ability of pollen types to remain airborne. In addition, for some locations, pollen deposition will also be influenced by the activities of animals, birds, or humans using the site area. Once deposited, other factors influence the eventual loss or recovery of specific pollen types. These factors include: 1) pollen recycling and/or mixture due to wind, water, humans, or burrowing animals; 2) the chemical composition of a pollen grain's wall [exine]; 3) the morphological shape and surface ornamentation type of each pollen type; and 4) the susceptibility of each pollen type to various types of degradation processes including those from mechanical, chemical, or biological agents (Bryant 1978, 1988; Bryant et al. 1994; Bryant and Hall 1993; Havinga 1964, 1984; Holloway 1989; King et al. 1975; O'Rourke 1990).

As noted by Jackson and Lyford (1999) and others, there are substantial differences among plant taxa in terms of their pollen production, methods of dispersal, and in the ability of their pollen grains to remain aloft and travel various distances from their dispersal source. These differences create an uneven distributional relationship between the amount of pollen that will fall to the ground (pollen rain) and the actual vegetational coverage of each plant taxon. Adding to this problem is the knowledge that a large number of plants produce small amounts of pollen that rarely is cast adrift into the atmosphere because the plants rely on insect or animal pollinators. These pollen types are seldom found in the normal pollen rain of a region even though the plants that produce them may compose a major portion of the vegetational coverage. Finally, animals and humans gathering plants for food or processing them for other purposes can artificially introduce additional amounts and types of pollen into the pollen rain of a specific location or region. For archaeologists, those "pollen additions," which are frequently found in the deposits of archaeological sites, can become useful clues used to infer various types of cultural activities.

Once deposited, pollen is subjected to a host of potential factors that will determine whether they will remain preserved over time. Another problem is that not all pollen types are created equal; therefore, some types will succumb to destruction much more rapidly than will other more durable types. That type of differential pollen preservation can create patterns of fossil pollen recovery in the sedimentary record, which might be easily misinterpreted by those with minimal knowledge of botany or palynology (Bryant et al. 1993; Reinhard et al. 2007).

One of the first agents that can affect pollen grains is mechanical degradation of the exine (pollen wall). After pollen is released from its source, it can become abraded or broken during the transportation phase before it falls to the ground and becomes part of the pollen rain. These alterations can result from impact or from changes in climatic conditions. Studies by Duhoux (1982), for example, have shown that atmospheric changes from low to high humidity can result in large numbers of exine ruptures in thin-walled pollen taxa such as those produced by plants in the Cupressaceae (Juniperus, and Thuja). Later, after being deposited, many of the thin-walled pollen types, as well as other pollen types, can become further abraded by various activities such as natural events (hurricanes, earthquakes, tornadoes) and animal disturbances (burrowing). The cultural practices of humans can also affect the preservation of deposited pollen. For example human events such as grassland burning, land surface modifications (temple building, irrigation systems, roadways, etc.), construction activities (house and temple building, pueblos, retaining walls, animal pens, etc.), and agricultural practices (slash-and-burn land clearing, plowing, surface flooding, etc.) will each alter the pollen rain and pollen deposition in different ways.

Mechanical abrasion of pollen can also occur during its dispersion cycle by various events including impact against objects, exposure to water, recycling, wind erosion, changes in temperature, changes in atmospheric or soil moisture levels, volcanic eruptions, and soil movement caused by earthquakes, creep, faulting, or uplifts.

The morphological structure and ornamentation of pollen walls seem to be important factors in determining their potential susceptibility to mechanical degradation. For example, protruding structures such as the bladders attached to the pollen grains of many conifer species (pine, spruce, fir), or the spines on the surface of some pollen types (composites, cotton, mallows, morning glory, etc.) have a tendency to break or erode through a variety of mechanical processes. In some cases, the actual appearance of a pollen grain may become so altered after the loss of an appended structure, or structures, that accurate identification is no longer possible. In addition, structural alteration by mechanical processes can cause cracks or exine weakening, thereby hastening the eventual destruction of the entire grain through other processes. An analogy would be the difference between a whole egg and one that had a hairline crack. The whole egg is much stronger and more durable than the one with even a tiny crack!

Soil chemistry, acting on the natural chemical composition of a pollen grain's exine is another factor that often plays an important role in determining pollen preservation. Although

the exine is composed mostly of cellulose and various types of proteins, there are also interlocking strands of a highly durable, polymer type material called sporopollenin. Early studies by chemists and palynologists including Brooks and Shaw (1968), Shaw (1971), Rowley and Prijanto (1977), and Rowley et al. (1990) have discovered a critical link between sporopollenin and the ability of a pollen type to become preserved. They found that differences in the amount of sporopollenin and differences in the specific molecular structure of variations of sporopollenin within the pollen exine enable some taxa to become preserved while other taxa with different variations and amounts of sporopollenin undergo rapid deterioration.

One of the primary indications of potential pollen preservation in sediments can be gained by determining the soil pH. By itself, pH is not entirely responsible for pollen destruction, but it is an important factor. As early as the 1950s, Dimpleby (1957) searched for causes of pollen deterioration in various types of soils. His experiments and research were the first to chart differences in pollen preservation caused by soil chemistry. His research revealed that most soils with an acidic pH seem to provide ideal conditions for pollen preservation. However, he noted that once acidic soils reach a pH level of being only weakly acidic (around 6.0), then significant pollen destruction begins to occur. Dimpleby even cautioned that attempts to recover fossil pollen from any soils with a pH greater than 6.0 would most probably result in failure. Since Dimpleby's original study, other studies by palynologists including Bryant et al. (1994), Hall (1981), and Martin (1963) have demonstrated that fossil pollen can be recovered in some alkaline soils with a pH as high as 8.9. Nevertheless, as noted by Bryant and Hall (1993), in most cases fossil pollen recovered from alkaline sediments is poorly preserved, is highly deteriorated, and frequently documents evidence of differential taxa preservation (i.e., many of the fragile pollen types have disappeared leaving behind only the most durable pollen types).

Related to Dimpleby's (1957) initial study on soil pH is Tschudy's (1969) later research on the effects of Eh (oxidation potential) on pollen preservation. Tschudy noted that Eh seems to be a better indicator for potential pollen preservation than pH. Sediments with a low Eh (from -1 to 0) represent a reducing, anaerobic type of condition where carbon dioxide and hydrogen sulfide are often present from the by-products of microbial respiration. This combination decreases the levels of oxygen and lowers the pH values. Thus, the creation of a negative Eh value generally results in the formation of a strongly reducing environment (Tschudy 1969). Because a reducing environment retards oxygen retention, which plays a key role in oxidizing organic compounds, it is ideal for pollen preservation. In addition,

reducing environments present unfavorable habitats for most microbes such as bacteria and fungi, some of which are known to feed on pollen. Therefore, a low Eh in a soil or sediment environment is ideal for pollen preservation.

One of the common types of sediments with a low Eh potential is an acidic peat bog, which is known to be one of the best places to recover fossil pollen. Likewise, as the Eh potential of sediments rises from 0 toward +1, it indicates oxidizing conditions that will speed the destruction of pollen. Oxidizing conditions create two types of destruction for fossil pollen. First, direct oxidation occurs when pollen grains come in contact with free oxygen. Second, indirect oxidation occurs when pollen grains come in contact with oxygenated water that percolates downward into sub-surface levels from the surface. This second type of oxidation is often the most destructive factor for fossil pollen and is especially common in the sediments of shallow ponds and lakes found in temperate environments or in well-drained sandy loam soils. The oxygenation of sub-surface soil and shallow lake deposits also provides an ideal habitat for species of pollen-eating bacteria and fungi.

Not all pollen types are created equal. The chemical composition of the pollen walls of some plant species is not nearly as durable as they are in other taxa. In addition, the structural morphology of the pollen wall plays an important role in determining whether a specific type of pollen grain will remain preserved in various types of sediments. In a 20-year study beginning in 1964 and ending in 1984, Havinga (1964, 1984) reported that the relationship between the morphological structure and percentage of sporopollenin in the wall composition of pollen grains seems to determine their susceptibility to eventual destruction by oxidation. He found, for example, that pollen grains having higher ratios of sporopollenin/cellulose in their walls tended to remain preserved much longer, even in soils with high pH and high Eh values. Likewise, those pollen taxa with thin walls and walls containing minimal amounts of sporopollenin tended to decompose very quickly in soils with either or both high pH and high Eh values. Subsequent to Havinga's 20-year study, additional research by Rowley et al. (1990) produced detailed scanning (SEM) and transmission (TEM) electron microscopy studies of how various types of decomposition occur in the pollen walls of different pollen taxa. The SEM and TEM studies by Rowley and his colleagues provide some of the most detailed images and explanations yet published outlining the destructive processes that affect pollen in various types of soil environments.

Biological agents including certain species of fungi and bacteria can cause damage to pollen grains that will speed their eventual destruction. Studies by Holloway (1981,

1989) noted that some species of Phycomycete fungi will seek out and feed on the nutrient materials in the cytoplasm of recently-deposited pollen grains. His experimental studies show that the filamentous threads of fungi, called hyphae, will often enter a pollen grain through one of the grain's natural aperture openings. Nevertheless, at other times the fungal hyphae seem to have the ability to dissolve areas of the pollen wall in order to provide an opening into the pollen grain. Both types of fungal attacks will weaken the wall structure of pollen and speed the grain's eventual destruction by other forms of chemical and mechanical degradation.

Some years earlier, Goldstein (1960) conducted experiments with various species of Phycomycete fungi and found that they were a causative factor in the destruction of fossil pollen. His original study revealed that certain species of Phycomycetes are selective in their preference for pollen types. One type of Phycomycetes, for example, would infect and feed only on certain types of conifer pollen grains even when a variety of other pollen types were more plentiful and more available in the substrate. That type of selective pollen destruction by certain species of fungi becomes an important factor to consider when trying to estimate the potential impact of differential pollen loss in certain sediments. Unlike Holloway's later study in 1989, Goldstein did not focus on how fungi actually damaged pollen grains. Instead, Goldstein was primarily concerned only with the percentage of pollen grains that were potentially infected and which pollen taxa were the most susceptible to fungal attack.

There were a number of fungal spores found in the samples I examined for this study. Nevertheless, I am not certain whether or not any of these may have been a contributing factor leading to the destruction and damage of fossil pollen in those samples.

Elsik (1966) was the first to note that bacterial degradation of pollen grains occurs. He found that certain bacteria, especially species of Actinomycetes, have an ability to degrade pollen. He found that even though much of the bacterial infection of pollen seems to occur soon after pollen grains are dispersed and when the pollen still contains cytoplasm, additional bacterial destruction continues to occur long after the pollen grains have lost their cytoplasm and have become part of the fossil record.

Finally, one of the most destructive forces for fossil pollen in sediments seems to come from repeated cycles of wetting and drying (Campbell and Campbell 1994). The walls of many pollen grains are elastic, which enable the fresh, viable pollen grains to expand and contract without destruction. Depending upon the changing levels of atmospheric humidity

some pollen grains may expand and contract dozens of times before they either reach their intended destination and begin the fertilization process, or fall to the surface as part of the pollen rain. For those that fall to the surface, they continue to expand and contract as moisture levels change. Eventually, this process weakens the grain and causes it to rupture or crack. Once weakened, pollen grains are much more susceptible to other processes of mechanical destruction.

Discussion

In their article Bryant and Hall (1993) discuss the validity aspects of fossil pollen samples and note that fossil pollen records should be suspect when they contain all three of the following: 1) fossil pollen from only a few plant taxa are represented, especially when the most abundant pollen taxa are from the most durable types of pollen with very distinctive and easy-to-recognize morphological features; 2) when a sample contains a high percentage of degraded and indeterminable pollen grains beyond recognition, and 3) when fossil pollen concentration levels are below 1,000 grains/gram or 2,500 grains/cubic centimeter of sediment. Two of those three aspects are present in all of the samples I examined for this project and all three aspects are present in 35% (8/23) of the total number of sediment samples I examined.

I found only 11 different fossil pollen taxa during the study of these samples (Tables B-2 through B-7). Not all 11 of those pollen taxa were found in any one of the individual samples. Because we do not have any control samples, it is difficult to guess the number of "expected" pollen taxa in that region of Texas. However, when working in the desert regions of the Big Bend National Park in West Texas we did compare the expected numbers of pollen taxa to the actual number of fossil pollen taxa recovered (Bryant et al. 1994). In that study we found that each of the fossil pollen samples contained approximately one-third or less of the expected diversity of pollen types normally found in the 89 surface samples examined from those same desert habitats. I would expect a higher and more diverse flora in the vegetation of Nueces County than is found in the desert regions of the Big Bend; therefore, control surface samples collected near the core sampling location should contain a wide diversity of pollen types. If so, then the 11 total fossil taxa recovered during this study may reflect, as it did in the Big Bend region, only one-third or less of the total number of original pollen taxa present in these deposits. This suggests, as it did for the Big Bend region, that we are recovering pollen evidence from only one-third or less of the original plant types that contributed pollen to these deposits. Therefore, like the Big Bend study, the recovered pollen record from the Oso Bay study has little validity in terms of presenting a true reflection

Table B-2. Fossil Pollen Recovered from Core Samples 88, 89, and 90

Oso Bay Samples Pollen Taxa	88		89		90	
	Total	%	Total	%	Total	%
APIACEAE (umbels)	0	0.0%	0	0.0%	1	0.5%
ASTERACEAE (sunflower-type)	40	19.7%	26	12.7%	12	5.8%
ASTERACEAE (ragweed type)	16	7.9%	11	5.4%	14	6.8%
Cheno-Ams (goosefoot)	38	18.7%	36	17.6%	52	25.2%
ONAGRACEAE (evening primrose)	0	0.0%	0	0.0%	0	0.0%
<i>Opuntia</i> (prickly pear cactus)	0	0.0%	0	0.0%	0	0.0%
<i>Pinus</i> (diploxylon pine)	4	2.0%	7	3.4%	3	1.5%
POACEAE (grass)	24	11.8%	32	15.6%	24	11.7%
<i>Polygonum</i> (knotweed)	0	0.0%	3	1.5%	3	1.5%
			0		0	
UNKNOWN	0	0.0%	0	0.0%	0	0.0%
			0		0	
<i>Degraded pollen beyond recognition</i>	81	39.9%	90	43.9%	98	47.6%
TOTAL POLLEN COUNTED	203	100.0%	205	100.0%	206	100.0%
Lycopodium Tracer Spores Counted	10		6		11	
Pollen Concentration/gram	27,405		46,125		25,281	

Table B-3. Fossil Pollen Recovered from Core Samples 91, 92, and 93

Oso Bay Samples Pollen Taxa	91		92		93	
	Total	%	Total	%	Total	%
<i>Alnus</i> (alder)	0	0.0%	0	0.0%	0	0.0%
ASTERACEAE (sunflower-type)	15	7.1%	11	5.3%	16	7.5%
ASTERACEAE (ragweed type)	11	5.2%	15	7.2%	22	10.4%
Cheno-Ams (goosefoot)	42	20.0%	31	14.8%	47	22.2%
<i>Ephedra</i> (joint fir)	0	0.0%	0	0.0%	0	0.0%
ONAGRACEAE (evening primrose)	0	0.0%	0	0.0%	0	0.0%
<i>Opuntia</i> (prickly pear cactus)	0	0.0%	0	0.0%	0	0.0%
<i>Pinus</i> (diploxylon pine)	4	1.9%	3	1.4%	2	0.9%
POACEAE (grass)	35	16.7%	44	21.1%	24	11.3%
<i>Polygonum</i> (knotweed)	2	1.0%	4	1.9%	3	1.4%
		0.0%				
UNKNOWN	0	0.0%	0	0.0%	0	0.0%
<i>Degraded pollen beyond recognition</i>	101	48.1%	101	48.3%	98	46.2%
TOTAL POLLEN COUNTED	210	100.0%	209	100.0%	212	100.0%
Lycopodium Tracer Spores Counted	14		12		6	
Pollen Concentration/gram	20,250		47,700		21,021	

Table B-4. Fossil Pollen Recovered from Core Samples 94, 95, and 96

Oso Bay Samples Pollen Taxa	94		95		96	
	Total	%	Total	%	Total	%
<i>Alnus</i> (alder)	0	0.0%	2	0.9%	0	0
ASTERACEAE (sunflower-type)	18	8.3%	21	9.9%	23	11.4%
ASTERACEAE (ragweed type)	21	9.7%	12	5.6%	18	8.9%
Cheno-Ams (goosefoot)	49	22.6%	54	25.4%	31	15.3%
<i>Ephedra</i> (joint fir)	0	0.0%	1	0.5%	0	0.0%
ONAGRACEAE (evening primrose)	0	0.0%	1	0.5%	0	0.0%
<i>Opuntia</i> (prickly pear cactus)	0	0.0%	2	0.9%	1	0.5%
<i>Pinus</i> (diploxylon pine)	5	2.3%	6	2.8%	2	1.0%
POACEAE (grass)	31	14.3%	33	15.5%	44	21.8%
<i>Polygonum</i> (knotweed)	4	1.8%	6	2.8%	2	1.0%
UNKNOWN	0	0.0%	0	0.0%	0	0.0%
<i>Degraded pollen beyond recognition</i>	89	41.0%	77	36.2%	81	40.1%
TOTAL POLLEN COUNTED	217	100.0%	213	100.0%	202	100.0%
Lycopodium Tracer Spores Counted	14		15		16	
Pollen Concentration/gram	21,021		19,170		17,043	

Table B-5. Fossil Pollen Recovered from Core Samples 97, 98, and 99

Oso Bay Samples Pollen Taxa	97		98		99	
	Total	%	Total	%	Total	%
<i>Alnus</i> (alder)	0	0.0%	0	0.0%	0	0.0%
ASTERACEAE (sunflower-type)	12	5.9%	16	7.3%	37	16.1%
ASTERACEAE (ragweed type)	14	6.9%	22	10.1%	30	13.0%
Cheno-Ams (goosefoot)	41	20.1%	51	23.4%	66	28.7%
<i>Ephedra</i> (joint fir)	0	0.0%	0	0.0%	0	0.0%
ONAGRACEAE (evening primrose)	0	0.0%	0	0.0%	0	0.0%
<i>Opuntia</i> (prickly pear cactus)	0	0.0%	1	0.5%	0	0.0%
<i>Pinus</i> (diploxylon pine)	4	2.0%	2	0.9%	2	0.9%
POACEAE (grass)	28	13.7%	44	20.2%	34	14.8%
<i>Polygonum</i> (knotweed)	6	2.9%	2	0.9%	1	0.4%
<i>Ulmus</i> (elm)	0	0.0%	0	0.0%	1	0.4%
UNKNOWN	0	0.0%	0	0.0%	0	0.0%
<i>Degraded pollen beyond recognition</i>	99	48.5%	80	36.7%	59	25.7%
TOTAL POLLEN COUNTED	204	100.0%	218	100.0%	230	100.0%
Lycopodium Tracer Spores Counted	10		12		14	
Pollen Concentration/gram	27,540		24,525		22,178	

Table B-6. Fossil Pollen Recovered from Core Samples 100, 101, and 102

Oso Bay Samples Pollen Taxa	100		101		102	
	Total	%	Total	%	Total	%
<i>Alnus</i> (alder)	0	0.0%	0	0.0%	0	0.0%
ASTERACEAE (sunflower-type)	29	13.1%	39	17.6%	29	13.2%
ASTERACEAE (ragweed type)	19	8.6%	18	8.1%	19	8.7%
Cheno-Ams (goosefoot)	62	28.1%	55	24.8%	29	13.2%
<i>Ephedra</i> (joint fir)	0	0.0%	1	0.5%	0	0.0%
ONAGRACEAE (evening primrose)	0	0.0%	0	0.0%	0	0.0%
<i>Opuntia</i> (prickly pear cactus)	1	0.5%	0	0.0%	0	0.0%
<i>Pinus</i> (diploxylon pine)	0	0.0%	3	1.4%	1	0.5%
POACEAE (grass)	65	29.4%	29	13.1%	39	17.8%
<i>Polygonum</i> (knotweed)	1	0.5%	0	0.0%	1	0.5%
<i>Ulmus</i> (elm)	0	0.0%	0	0.0%	0	0.0%
UNKNOWN	0	0.0%	0	0.0%	0	0.0%
	0	0.0%	0	0.0%	0	0.0%
<i>Degraded pollen beyond recognition</i>	44	19.9%	77	34.7%	101	46.1%
TOTAL POLLEN COUNTED	221	100.0%	222	100.0%	219	100.0%
Lycopodium Tracer Spores Counted	16		36		80	
Pollen Concentration/gram	18,646		8,325		3,695	

Table B-7. Fossil Pollen Recovered from Core Samples 178, 179, 180, 181, 275, 276, 277, and 278

Oso Bay Samples Pollen Taxa	178	179	180	181	275	276	277	278
ASTERACEAE (sunflower-type)	0	0	0	0	0	2	2	0
ASTERACEAE (ragweed type)	0	0	0	0	0	0	1	0
ASTERACEAE (dandelion type)	0	0	0	0	1	0	0	0
Cheno-Ams (goosefoot)	1	0	0	0	3	2	0	0
<i>Pinus</i> (diploxylon pine)	2	0	0	0	1	0	0	0
POACEAE (grass)	0	0	0	0	2	2	3	0
<i>Quercus</i> (oak)	0	0	1	0	0	0	0	0
<i>Ulmus</i> (elm)	0	1	0	1	0	0	0	0
Lycopodium Tracer Spores Counted	50	50	50	50	50	50	50	50
Pollen concentration/gram	54	27	27	27	189	162	162	0

of the actual pollen record that was first deposited in these core samples as part of the regional pollen rain.

Other evidence of the severity of pollen destruction in the core samples examined for this project is the high numbers of degraded grains in each sample (Figure B-1). The percentage of degraded pollen, ranging from a low of 19% to a high of 48% in these samples represent pollen grains that are recognizable only as being pollen at one time but are now so altered that their identification is not possible. As first noted by Dean (1991) in most fossil pollen studies when the percentage of indeterminate and highly degraded pollen exceeds 10% in an individual sample, then that sample's pollen data should come under serious question and should not be considered an accurate reflection of the true pollen record.

A third line of evidence is generally the actual pollen concentration value of each sample. The pollen concentration values found in these 23 samples ranged from a low of 0 fossil pollen grains/gram of sediment to a high of 47,700 fossil pollen grains/gram of sediment. Under most circumstances fossil pollen concentration values near 50,000/gram in a sample would be considered excellent for interpretative purposes. However, in sample # 93, which contained the highest fossil pollen concentration value of in this study, 46% of those pollen grains were degraded beyond recognition as to which taxon they belonged. Thus, even with what is often considered an "outstanding" pollen concentration value, the total number of degraded pollen grains is nearly 50% and the total number of pollen taxa found in the sample was only five types. In circumstances, such as this one, the high fossil pollen concentration value is meaningless as an indication of preservation because so much of the original pollen is missing or is highly degraded. Control surface samples from this same region would have enabled us to determine an approximate modern pollen concentration value for the normal pollen rain at the coring locale.

Pollen Types

There are only 11 different pollen types represented in the fossil record of these samples. Of those types, the three dominant pollen taxa are types from the composites, Chenosporites, and grasses.

Asteraceae: There are hundreds of species of composites growing throughout the region of South Texas. Some of those species, such as those identified as belonging to the "ragweed type" produce vast numbers of wind-dispersed pollen. Other composite species produce flowers that are insect-pollinated. Sunflowers and other composites that are closely related to them produce flowers that are insect-pollinated. Those pollen

types are recorded in this study as belonging to the primary category termed "sunflower-type." Although ethnographic sources record that a number of composite species were selected and used as food by aboriginal cultures in Texas and throughout much of western North America (Casteretter 1935; Moerman 1998; Palmer 1887), a few species of composites get the most attention because they were often sought for their seeds, which are rich in fats and oil (Moerman 1998). Whether or not the presence of both ragweed and sunflower types of pollen in these core deposits reflect the use of these plants by some of the early cultures living at or near the Cayo del Oso Site (41NU2) is not known and cannot be proven by the pollen record from these core samples. Although both of these pollen types (Figure B-2) individually exceed 10% of the total pollen in many of these samples, this may not be an accurate record of the abundance of these plants in the ancient environment. In fossil deposits where there is a high degree of pollen degradation the high levels of composite pollen often is not an accurate reflection either of their economic use or the abundance of those plants in the local environment. Bryant and Hall (1993) note that all forms of composite pollen are among the "most durable" pollen types found in the Holocene fossil record and that those pollen grains are often identifiable even when most other fossil pollen taxa have deteriorated under unfavorable conditions of preservation. Therefore, as a general rule, when organic preservation is poor, composite pollen tends to become "overrepresented" in terms of its recovered presence, as opposed to its original presence in the ancient pollen rain of the sampled locale.

Cheno-Ams: The fossil pollen in this category of plants may have come from a wide variety of species many of which are drought resistant, or tolerance of saline environments, or grow quickly and flourish in disturbed soils. These are the types of plants that will often replace grasses, forbs, and even small bushes when environmental conditions become harsh or after vegetational changes occur as a result of fires, erosion, hurricanes, or encroachment from marine environments. Most of the plants in the Chenosporites category produce large quantities of small and very durable wind-pollinated pollen grains that are easily dispersed by the wind. The mature plants also produce thousands of small seeds, which were often gathered and eaten by many prehistoric and historic cultures living throughout the American Southwest and northern Mexico (Feweks 1896; Moerman 1998). Therefore, as with the records of Asteraceae pollen, it is tempting to infer that the parent plants producing the Chenosporites pollen were widespread near the Cayo del Oso Site and that perhaps ancient cultures collected and used the seeds. On the other hand, I personally believe that the high percentages of fossil Chenosporites pollen in these core samples (Figure B-3) result from the destruction and elimination of other less durable pollen types in these sediments. That type of condition, confirmed by the high percentage of degraded and

Scale for all photos: Space between numbers is equal to 25 microns.

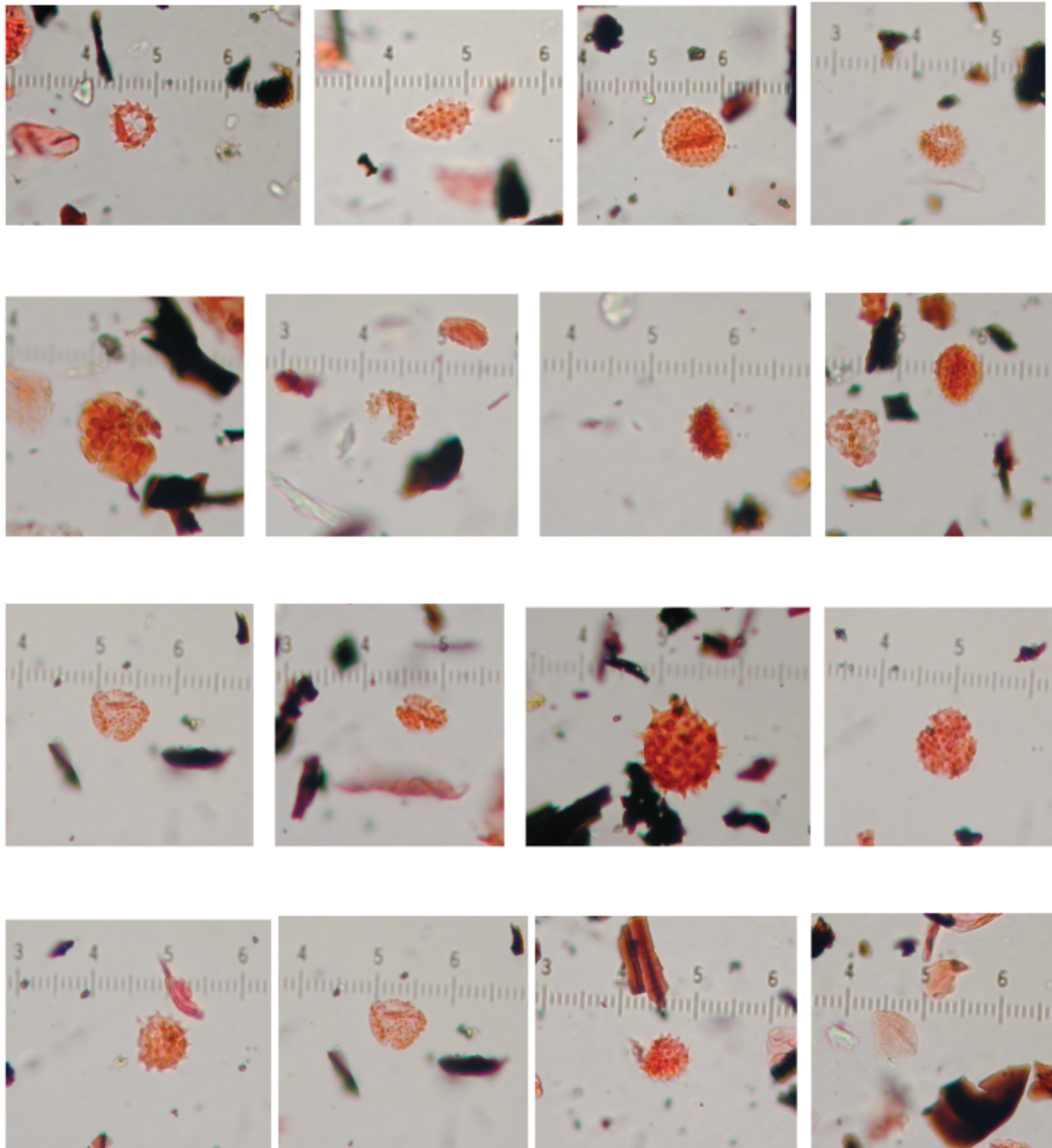


Figure B-2. Fossil Asteraceae Pollen Types Recovered from the Core Samples.

indeterminate fossil pollen in every sample examined, would enable Cheno-Ams fossil pollen to become over represented because they are durable types that remains identifiable even with badly degraded.

Poaceae: We know from ethnographic reports and from studies of human coprolites found throughout West Texas and the American Southwest (Bryant and Williams-Dean 1975; Chamberlin 1911; Moerman 1998; Stock 1983; Sobolik

1991) that Native Americans collected and ate various types of grass seeds including those from foxtail barley, wild-rye, and millets. Although it is possible that some of the fossil grass pollen found in the soils of these core samples may have come from species that could be considered economically important for ancient cultures, it is more probable that the pollen came from natural sources (Figure B-4). Grasses are a common plant growing around the edges of small ponds, such as the one sampled for these core samples, and along the shorelines of marine and brackish inlets. Grasses also flourish

Scale for all photos: Space between numbers is equal to 25 microns.

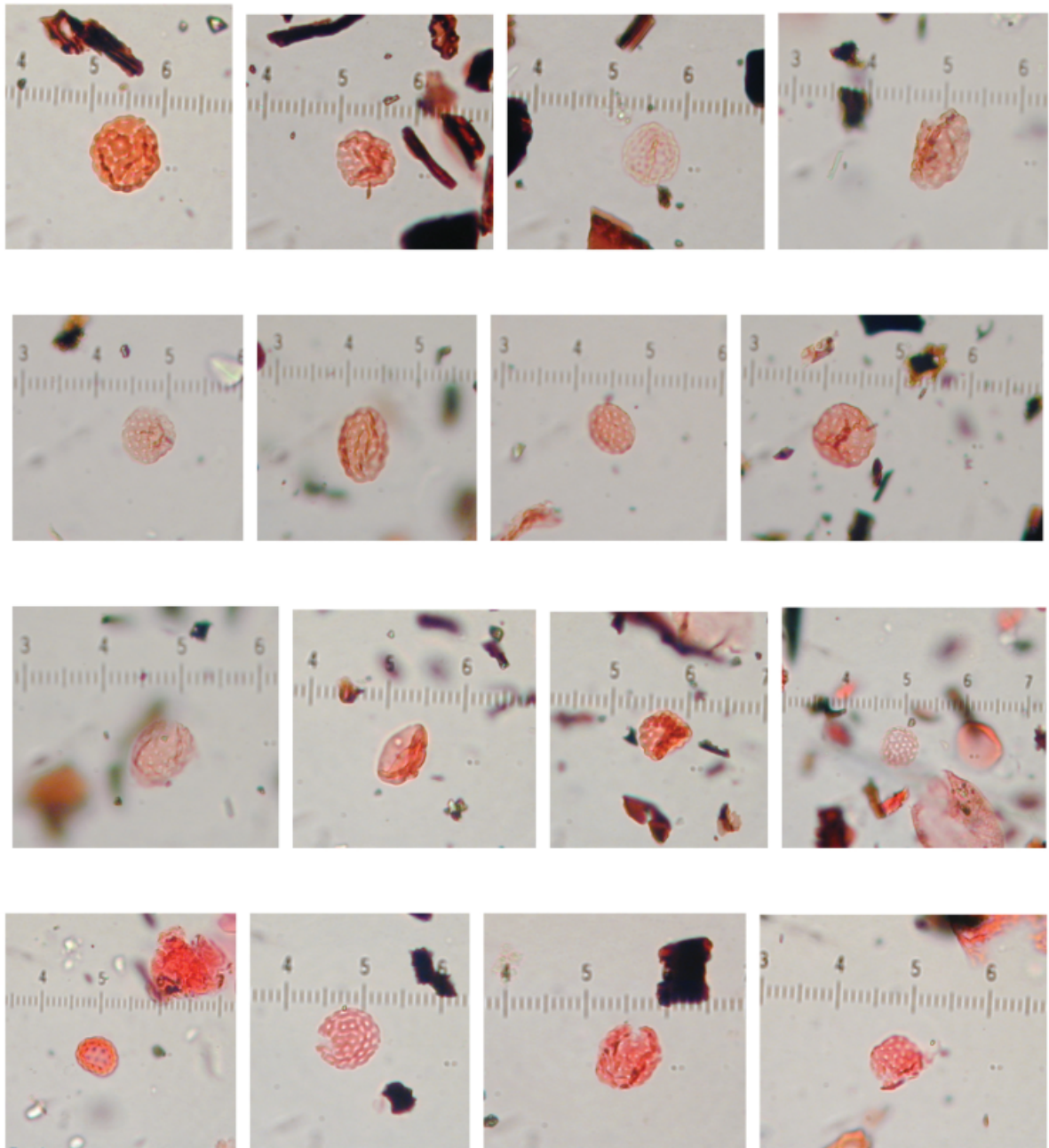


Figure B-3. Fossil *Cheno-Ams* Pollen Types Recovered from the Core Samples.

in coastal areas, such as the region around the Cayo del Oso Site. Finally, like the pollen from species of composites and the pollen from species of *Cheno-Ams*, grass pollen is fairly durable and is often recognizable in the fossil record even though the grains have become severely degraded, as noted by Bryant and Hall (1993).

Ephedra: Ephedra plants grow well in hot climates where water is at a premium. They can also grow at various

altitudes, which make them a common plant throughout many of the arid regions in the western portion of the United States. However, even though joint fir pollen was found in two of the 23 samples, I doubt any of that pollen came from local sources (Figure B-5). That suggests that the wind-dispersed grains of joint fir pollen probably came from some far distant source before being deposited in the pond where the core samples were collected. It is also important to remember that Bryant and Hall (1993) note

Scale for all photos: Space between numbers is equal to 25 microns.

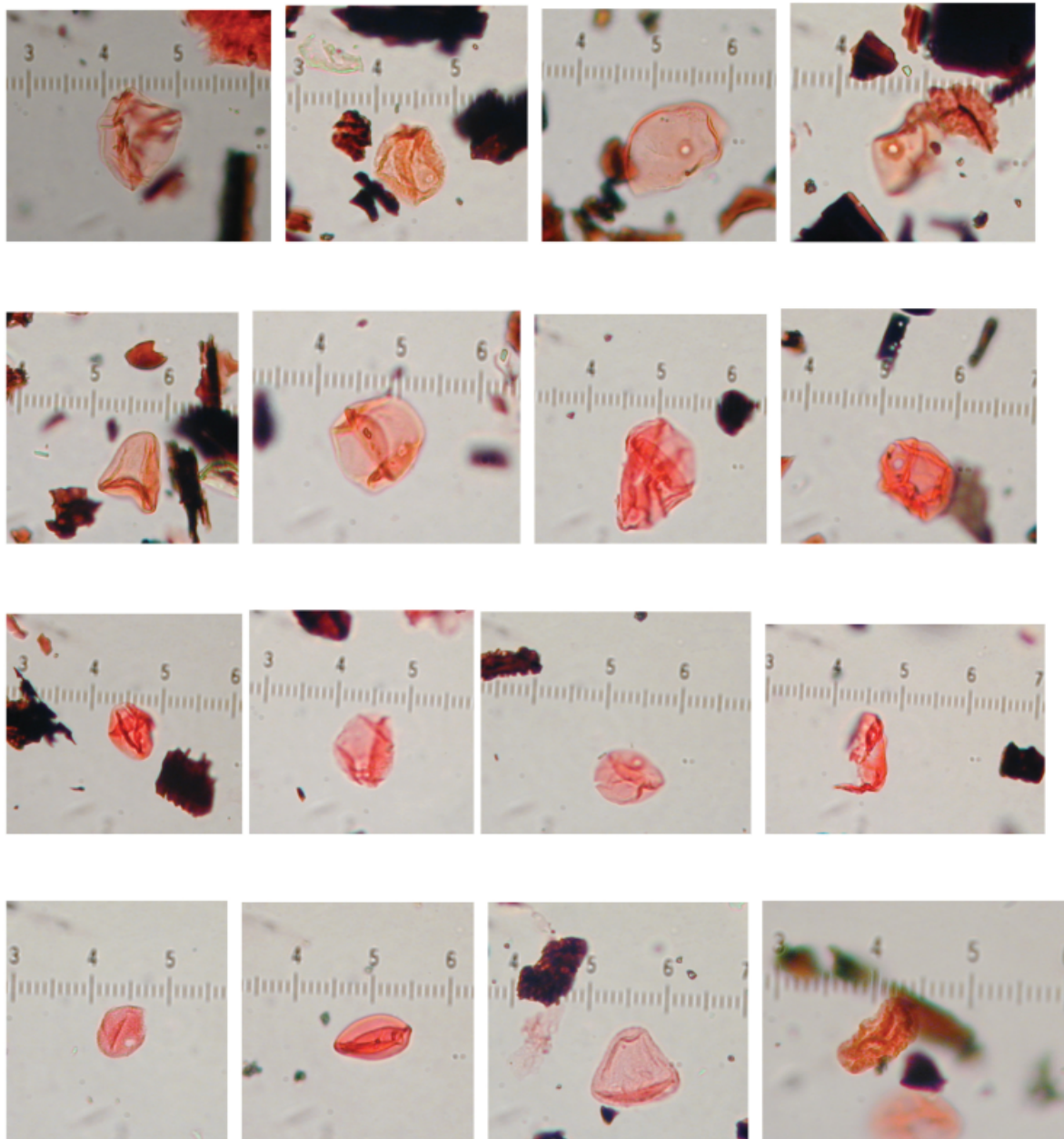


Figure B-4. *Fossil Poaceae Pollen Types Recovered from the Core Samples.*

that joint fir plants produce one of the most durable types of pollen grains and that it is able to survive in soils long after many fragile pollen types have become severely weathered or have completely disappeared.

Miscellaneous pollen types: There are a few other fossil pollen taxa represented in the pollen record of these core samples, some of which may reflect local parent plants (Figure B-5). Traces of prickly pear cactus (*Opuntia*) pollen were found in

4 of the 23 samples. Individual pollen grains from prickly pear cactus are very large and fragile. Even under “good” depositional conditions because of the fragile nature of prickly pear cactus pollen many of the grains become broken and fragmented. Once broken, many of the small pieces of prickly pear cactus pollen are still recognizable even though most of the grain has been degraded. All of the cactus pollen represented in the pollen counts of these core samples were fragments of grains rather than whole cactus pollen grains.

Scale for all photos: Space between numbers is equal to 25 microns.

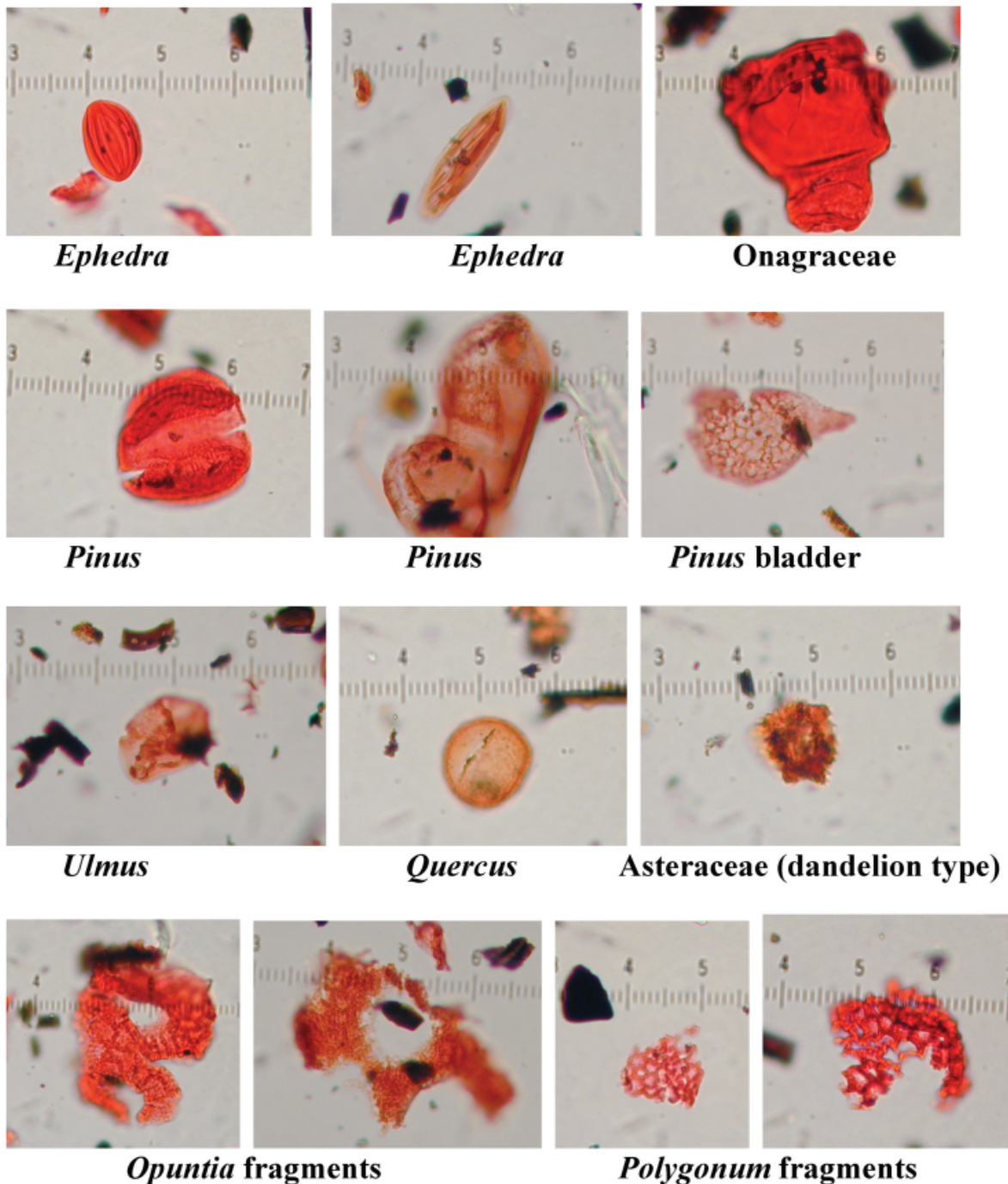


Figure B-5. Fossil Pollen Types from Various Taxa Recovered from the Core Samples.

There were also fragments of knotweed (*Polygonum*) pollen found in 13 of the 23 samples. Knotweeds are common plant that grow well in disturbed areas and are common in modern agricultural fields where they are considered a weed. These plants can tolerate many different conditions, but prefers moist soils. Today they are common throughout most regions

of the eastern and central North America. They are also very common throughout most areas of Texas. The pollen grains of some species of knotweed, such as those of *Polygonum longistylum*, have a fairly unique morphology, therefore even when the pollen grains have been broken into small fragments, many of the fragments can still be identified. The

majority of the knotweed pollen found in the 13 core samples consisted of fragments of pollen grains rather than the whole pollen grains.

These samples contained a few pollen grains from arboreal sources such as pines, oaks, and elms. Sixteen out of the 23 samples contained at least one pine pollen grain. Pines do not grow naturally in Nueces County, the location of the core samples. However, pine trees produce copious amounts of pollen that have air sacs and are designed to travel easily on air currents. It is not uncommon for a few pine pollen grains to travel hundreds of miles from their site of dispersion. Thus, I suspect that the few pine pollen grains found in these samples are the result of long distance transport rather than coming from trees growing locally. The single oak pollen grain found in sample 180 is most probably a contaminant rather than a pollen grain that was originally deposited in the sediments of that sample. Oak trees produce large quantities of pollen that are airborne and can travel long distances. Nevertheless, because the single oak pollen grain was the only pollen grain found in sample 180 and because it was in pristine condition, I

suspect that at some point it may have been incorporated into the sample, perhaps when the sub sample was being removed from the core for pollen studies. Single pollen grains of elm were found in three samples and may have come from local elm trees. Today elm trees are common in the vegetation of Nueces County and I suspect they were probably present in the region in prehistoric times as well.

Summary

Unfortunately, the recovered pollen from all of the core samples I examined reflect severe cases of pollen degradation and the probable loss of a large number of pollen taxa and pollen grains between the time the pollen rain was originally deposited and the time the samples were collected and examined for fossil pollen. When destruction of fossil pollen is as severe as it appears to be in these samples, little interpretative value can be gleaned from the resulting data. To use the pollen data otherwise would result in speculations based on unsupported assumptions.

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APPENDIX C:
Plant Remains from 41NU2

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The purpose of this analysis is to provide an assessment of the botanical assemblages from 41NU2, the Cayo del Oso site. A total of 66 flotation samples were submitted for analysis. The data will be utilized to assess the nature and condition of the plant remains from these sites and provide some evidence for plant utilization and local environmental conditions.

Methods

Flotation is a method of recovering organic remains from archeological sediments by using water to separate heavy or soluble inorganic particles from plant parts and small animal bone. The material floating to the surface is called the light fraction, and this is caught on a fine mesh screen or strainer. The material that sinks to the bottom is the heavy fraction and it is also caught on a fine mesh screen. Most of the soil including clay and silt is suspended in water and passes through the screens and is either recycled or discarded. In this study, the light fractions were submitted for analysis.

The analysis followed standard archeobotanical laboratory procedures. The volume of the light fraction is first measured. In most cases up to 100 ml of light fraction from each sample is set aside for analysis. Then the portion to be analyzed is passed through a nested set of screens of 4mm, 2mm, 1mm, and 0.450mm mesh and examined for charred material, which is separated for identification. The heavy fraction is also scanned for charcoal, which is set aside for identification. The results are combined with identifications from each corresponding light fraction. In the current study, charred plant material was so scarce that I sorted through the entire volume of each light fraction that was submitted for analysis.

Plant material is sorted into two categories—woody fragments, and seed/fruit fragments including maize or agave parts when present. Identification of carbonized wood was accomplished by using the snap technique, examining the fragments at 8 to 45 magnifications with a hand lens or a binocular dissecting microscope, and comparing the material to samples in the archeobotanical herbarium. All seed identifications were made using seed manuals and reference collections at Shumla Archeobotanical Services. Only charred plant material is included in the analysis, because uncarbonized material is consumed by insects, fungi and bacteria and does not survive more than a few years in the deposits of open sites.

Up to 25 wood charcoal fragments large enough to be manipulated are examined and identified from each flotation sample. Fragments smaller than 2- or 3- mm are difficult

to manipulate, and are usually placed in the indeterminate category. When a sample contains more than 25 fragments, the rest of the material is scanned to make sure that no other taxa are present. Then the volume of the charcoal is measured and included along with its weight in the report. Due to the scarcity of carbonized plant remains in the Cayo del Oso samples, all the charred material was examined carefully. The results for each sample are presented in tabular format.

Disturbance Indicators. Sample content may be affected by various biological disturbance factors, including insect or small mammal activity, and plant root growth. In an effort to assess this impact, the amounts of insect parts, termite pellets, gastropods, mammal remains (including fecal pellets), and modern uncharred seeds are estimated for each flotation sample. These amounts are reported on a scale of 1-5 (+), 6-25 (++), 26-50 (+++), and over 50 (++++). Termite pellets occur in higher numbers when samples are taken from an area containing wood that has been exposed to the elements for a long time before burning. In the desert, this can occur in dead trees or roots, in which case the termite pellets can appear in any locus that this wood is burned, such as in a hearth or roasting pit. However, evidence of termite infestations seems to be more frequent and intense in samples drawn from the remains of burned prehistoric habitations with vertical elements constructed of wood.

Results and Discussion

Results of the analysis are presented in Tables C-1 and C-2. Table C-1 lists the sample proveniences along with charred seed taxa counts and total charred seed counts from each sample. Identifications and counts of plant material recovered from each sample, including the macrobotanical sample, appear in Table C-2. Recovery of charred plant material in the flotation samples generally sparse. The total charcoal in each sample varied from 0 or <.01 g to 1.1 g. Disturbance indicators were abundant, and sand, usually absent from light fractions, was present in many samples. Roots and gastropods constituted the majority of the disturbance indicators, and several samples contained uncharred seeds. Some of the light fractions submitted for analysis were very small; in fact, 24 of the unsorted light fractions measured less than 10-ml, and four measured 1-ml or less.

Three samples contained seeds or fruits; seeds identified in the samples include cheno-am and prickly pear. Wood charcoal was widespread, in small quantities, throughout the samples, occurring in 36 samples. Thirty samples contained either charcoal flecks or no carbonized plant material.

Table C-1. Flotation Sample Summary

Flot Sample Number	FS #	Feature	Light fraction vol. (mL), wt. (g)	Roots (r), Insect Parts (ip), Gastropods (g)	Uncharred Seeds	Number of Charred Seed Taxa (Agave, Yucca)	Total Charred Seeds	Total Charcoal (g)
1	7	1	27; 1.9	r+++	0	0	0	0
2	48	9	17; 3.6	sand	0	0	0	<0.1
3	65	13	35; 7.7	r+++ (sand)	0	0	0	<0.1
4	73	14	10; 3.1	r+++	0	0	0	0.1
5	128	40	4; 0.1	sand	0	0	0	0
6	144	48	6; 0.7	r+++	0	0	0	0.2
7	147	51	30; 7.2	sand	0	0	0	0.2
8	149	53	8.2; 41	g+++ (sand)	0	0	0	<0.1
9	154	54	24; 6.5	r+++ g+	0	0	0	1.1
10	155	54	2; <.1	sand	0	0	0	0
11	175	61	12; 2.7	r+++	0	0	0	<0.1
12	199	73	31; 16.2	g+++ (sand)	0	0	0	0
13	209	75	33; 8.3	r++ g+++	1	0	0	<0.1
14	222	74	37; 2.2	r+ g++	0	0	0	0.2
15	227	78	15; 1.2	r+++ g+	0	0	0	<0.1
16	232	84	29; 7.4	r+++ g+	0	0	0	0
17	238	57	18; 1.6	r+++ g+	0	0	0	0
18	239	57	12; 3.6	r+++ g+	0	1	1	<0.1
19	240	88	12; .8	r+++	0	0	0	0
20	242	89	24; 0.8	r+++	0	0	0	0
21	246	90	6; 0.6	r+++ (sand)	0	0	0	0
22	249	92	17; 3.9	r+++ g+	0	0	0	0
23	251	93	12; .4	r+++	0	0	0	0.1
24	252	94	4; 0.8	r+++ g+	0	0	0	0
25	260	99	24; 4.1	r+++	0	0	0	<0.1
26	261	100	15.4; 97	r+++ g+	0	0	0	0.1
27	270	65	87; 9.3	r+++ g+	0	0	0	<0.1
28	275	102	54; 13.2	r+++ g+	0	0	0	<0.1
29	276	102	79; 16.2	r+++ g++	0	0	0	0.2
30	332	121	4; 3.1	g+	0	0	0	<0.1
31	333	121	6; .2	g+	0	0	0	<0.1
32	341	125	14; 7.1	r+++	0	0	0	<0.1
33	400	136	8; 0.9	r+++ (sand)	0	0	0	0
34	391	133	17; 0.7	r+++ g+	0	0	0	0
35	452	140	4; 0.2	r+++	0	0	0	0
36	457	141	3; 03	r+++ (sand)	0	0	0	0
37	434	139	19; 1.2	r+++ g+	0	0	0	0
38	462	144	0.8; <0.1	r+++	0	0	0	0
39	459	143	12; 0.6	r+++	0	0	0	<0.1
40	313	???	12; 1.6	r+++ g+	0	0	0	<0.1
41	319	114	100; 8.8	r+++ g+	0	0	0	0.5
42	321	117	100; 8.1	r+++ g+++	0	0	0	<.1

Table C-1. Continued...

Flot Sample Number	FS #	Feature	Light fraction vol. (mL), wt. (g)	Roots (r), Insect Parts (ip), Gastropods (g)	Uncharred Seeds	Number of Charred Seed Taxa (Agave, Yucca)	Total Charred Seeds	Total Charcoal (g)
43	474	147	97; 9.8	r+++ (sand)	4	0	0	0.1
44	485	149	6; 0.3	r+++	0	0	0	<0.1
45	488	150	17; 2.0	r+++ (sand)	0	0	0	<0.1
46	497	152	3; 0.2	r+++	0	0	0	0
47	494	151	27; 12.7	r+++; g++	0	0	0	0
48	505	155	14; 3.2	r+++; g+	4	0	0	0
49	507	156	19; 1.2	r+++	2	1	1	<0.1
50	510	157	3; 0.1	r+++	0	0	0	0
51	511	158	<0.1; <0.1	sand	0	0	0	0
52	514	159	20; 1.1	r+++	0	0	0	0
53	502	154	5; 0.7	r+++; g+	0	0	0	0
54	516	161	15; 1.3	r+++	0	0	0	0
55	519	162	2; <0.1	r+++	0	0	0	0
56	524	163	21; 1.3	r+++; l++	0	0	0	0
57	530	164	22; 2.2	r+++	0	0	0	<0.1
58	532	165	12; 2.1	r+++	1	0	0	0.2
59	551	170	4; .3	r+ (sand)	0	0	0	0
60	534	166	8; 1.3	r+++	0	0	0	0
61	536	167	7; 0.3	r+++; g+	0	0	0	<0.1
62	541	168	32; 4.1	r+++	0	0	0	0.2
63	546	169	1; 0.2	sand	0	0	0	0
64	248	92	17; 0.8	r+++	0	0	0	0.2
65	477	148	0.5; <0.1	r+	0	0	0	<0.1
66	499	153	7; 1.1	r+++; g+	0	0	0	0

Taxa Summaries: Wood

Five wood taxa, mesquite (*Prosopis glandulosa*), woody legume (*Prosopis* or *Acacia* spp.), lotebush (*Ziziphus obtusifolia*), oak (*Quercus* sp.), and brasil (*Condalia* sp.) were identified in the samples, as noted in Table C-2. Woody legume is a general wood type that includes both mesquite and several species of acacia. The wood anatomy of these trees or shrubs is similar and often small fragments of charcoal are difficult to distinguish. In addition to the identified material several very small fragments of wood were placed in the indeterminate category. Two of these wood types, oak and mesquite, have been recorded in archeological sites and ethnographic or historic records. Brasil and lotebush were likely used for wood, and lotebush produces a small, edible fruit, but we have no ethnographic mention of their use by Native Americans.

These taxa are typical components of Taumalipan Thornscrub vegetation which dominates the clay dunes and sand sheets of

the Texas and Mexican coastal plain from the Coastal Bend to southern Tamaulipas, Mexico.

Mesquite and Woody Legume. Carbonized mesquite wood fragments were identified in five features and woody legume in 14 features. Unfortunately, no seeds or pod fragments were found in samples from 41NU2. Very few flotation samples have been examined from sites located in the Coastal Bend or southern Texas. Perhaps for this reason, archeologists have not recovered much evidence for the use of mesquite as food in southern Texas. However, mesquite seeds have been identified from three sites in Maverick County near Eagle Pass, 41MV136, 41MV160, and 41MV164 (Dering 2005). Mesquite charcoal has been identified from the Hinojosa site in Jim Wells County (Jones 1986). Mesquite wood charcoal also has been identified from several sites in southern Texas, including 41ZP364 in Zapata County and the Lino site in Webb County, (Dering 2000a and b).

Table C-2. Plant Remains Ordered by Feature Number

Flot #	FS #	Feature	Taxon	Common	Part	Count	Wt (g)
1	7	1	No carbonized plant remains				
2	48	9	Indeterminate	NA	Wood	1	<0.1
3	65	13	Fabaceae	Woody legume	Wood	2	<0.1
4	73	14	Fabaceae	Woody legume	Wood	10	0.1
5	128	40	No carbonized plant remains				
6	144	48	Fabaceae	Woody legume	Wood	11	0.2
7	147	51	Fabaceae	Woody legume	Wood	14	0.2
8	149	53	Indeterminate	NA	Wood	3	<0.1
9	154	54	Fabaceae	Woody legume	Wood	24	1.1
10	155	54	No carbonized plant remains				
11	175	61	Indeterminate	NA	Wood	1	<0.1
12	199	73	No carbonized plant remains				
13	209	75	Cheno-am	Chenopodium or Amaranthus	Seed (fruit)	1	<0.1
13	209	75	Fabaceae	Woody legume	Wood	5	<0.1
13	209	75	Indeterminate	NA	Wood	24	<0.1
14	222	74	Indeterminate	NA	Wood	22	<0.1
14	222	74	Quercus sp.	Oak	Wood	9	0.1
15	227	78	Fabaceae	Woody legume	Wood	3	<0.1
16	232	84	No carbonized plant remains				
17	238	57	No carbonized plant remains				
18	239	57	Cheno-am	Chenopodium or Amaranthus	Seed (fruit)	1	<0.1
18	239	57	Indeterminate	NA	Wood	5	<0.1
19	240	88	No carbonized plant remains				
20	242	89	No carbonized plant remains				
21	246	90	No carbonized plant remains				
22	249	92	No carbonized plant remains				
23	251	93	Indeterminate	NA	Wood	5	0.1
24	252	94	No carbonized plant remains				
25	260	99	Fabaceae	Woody legume	Wood	6	<0.1
25	260	99	Indeterminate	????????????	????????	1	<0.1
26	261	100	Indeterminate	NA	Wood	13	0.1
27	270	65	Indeterminate	NA	????????		<0.1
28	275	102	Indeterminate	NA	Wood	8	<0.1
28	275	102	Zea mays ?	Maize	Cupule?	1	<0.1
29	276	102	Indeterminate	NA	Wood	6	<0.1
29	276	102	Prosopis glandulosa	Mesquite	Wood	11	0.2
29	276	102	Ziziphus obtusifolia	Lotebush	Wood	3	<0.1
30	332	121	Indeterminate	NA	Wood	7	<0.1
31	333	121	Fabaceae	Woody legume	Wood	8	<0.1
32	341	125	Fabaceae	Woody legume	Wood	11	<0.1
33	400	136	No carbonized plant remains				
34	391	133	No carbonized plant remains				
35	452	140	No carbonized plant remains				
36	457	141	No carbonized plant remains				
37	434	139	No carbonized plant remains				

Table C-2. Continued...

Flot #	FS #	Feature	Taxon	Common	Part	Count	Wt (g)
38	462	144	No carbonized plant remains				
39	459	143	Fabaceae	Woody legume	Wood	1	<0.1
40	313	???	Indeterminate	NA	Wood	6	<0.1
41	319	114	<i>Prosopis glandulosa</i>	Mesquite	Wood	14	0.5
42	321	117	Fabaceae	Woody legume	Wood	5	<0.1
42	321	117	Indeterminate	NA	Wood	18	<0.1
43	474	147	Indeterminate	NA	Wood		0.1
44	485	149	Indeterminate	NA	Wood	1	<0.1
45	488	150	Indeterminate	NA	Wood	2	<0.1
46	497	152	No carbonized plant remains				
47	494	151	No carbonized plant remains				
48	505	155	No carbonized plant remains				
49	507	156	<i>Opuntia</i> sp.	Prickly Pear	Seed	1	<0.1
49	507	156	<i>Prosopis glandulosa</i>	Mesquite	Wood	5	<0.1
50	510	157	No carbonized plant remains				
51	511	158	No carbonized plant remains				
52	514	159	No carbonized plant remains				
53	502	154	No carbonized plant remains				
55	519	162	No carbonized plant remains				
56	524	163	No carbonized plant remains				
57	530	164	Indeterminate	NA	Indeterminate	4	<0.1
58	532	165	<i>Prosopis glandulosa</i>	Mesquite	Wood	22	0.2
58	532	165	<i>Condalia</i> sp.	Brasil	Wood	3	<0.1
59	551	170	No carbonized plant remains				
60	534	166	No carbonized plant remains				
61	536	167	Fabaceae	Woody legume	Wood	2	<0.1
62	541	168	<i>Prosopis glandulosa</i>	Mesquite	Wood	13	0.2
63	546	169	No carbonized plant remains				
64	248	92	Fabaceae	Woody legume	Wood	6	0.1
64	248	92	Indeterminate	NA	Wood	9	0.1
65	477	148	Indeterminate	NA	Wood	2	<0.1
66	499	153	No carbonized plant remains				

Throughout its distribution, mesquite was a vital resource for food, fuel, medicine, and implement-making. Cabeza de Vaca, traveling through either southern Texas or northern Mexico in the 1520's, is the first European to observe and record the use of mesquite by Native Americans. While he was living among the Cuchendados, he observed how they prepared mesquite pods for food. They dug a pit and filled it with pods, and then pounded them with a wooden pestle as thick as a man's thigh. They mixed the pod meal with handfuls of earth and consumed it raw, discarding the seeds and woody casings (Campbell and Campbell 1981; Krieger 2002).

Oak. Oak wood was identified in single sample from Feature 74. Oaks grow in small, densely packed mottes on sandy

soils in the region. The most common of these, *Quercus fusiformis*, also grows on the Edwards Plateau and adjacent South Texas Plains.

Cabeza de Vaca did not mention the use of oak or acorns in his account, however, oak wood is a very common component in archeological sites on and near the Edwards Plateau. In samples I analyzed from 41BX100 a small heating feature contained over 50 charred acorn cotyledons (Dering 2001). The presence of oak in the Cayo del Oso samples suggests that it was used for firewood, and it is possible that acorns also were harvested in the area. However, because of the small size of the oak mottes and abundance of other food resources, oak probably was not a staple or even a preferred food.

Brasil and Lotebush. Both brasil (*Condalia* sp.) and lotebush (*Ziziphus obtusifolia*) were identified in the Cayo del Oso samples. Some species of *Condalia* (brasil) produce small edible fruit, and the wood provides an adequate source of fuel. Lotebush produces a small edible fruit that was utilized by Native Americans across the greater Southwest.

The small wood fragments recovered from Feature 165 were likely used as fuel in a small heating feature. They constitute the first identification of brasil from an archeological site in southern Texas. However, *Condalia* seeds were noted in 3,000 year-old deposits in the well-preserved rockshelter deposits of Hinds Cave in the Lower Pecos, located on the far western edge of the South Texas Plains (Dering 1979). The Maricopa, Mohave, and Pima gathered lotebush fruits and ate them raw or boiled them into a syrup (Castetter and Bell 1951; Curtin 1984; Russell 1908).

Seeds and Other Plant Parts

Cheno-am. Many plants of the two genera *Chenopodium* (goosefoot) and *Amaranthus* (pigweed or amaranth) produce seeds that are quite similar and difficult to separate, especially when carbonized or partially decomposed. They also have similar growth habits and economic uses. When seeds are recovered that could be identified to either genus, the term “cheno-am” is used (cf. Adams 1994:177). Therefore, Cheno-am refers to the seed-like fruit of either goosefoot (*Chenopodium*) or pigweed (*Amaranthus*). When the fruit has been modified by charring it is often difficult or impossible to identify to the genus level, and it is placed in the artificial category cheno-am. Both plants were utilized in a similar fashion, and both contributed fruit and greens to the diet of Native Americans throughout the greater Southwest. Several charred seeds identified as *Chenopodium berlandieri* were recovered from Late Prehistoric contexts at the Hinojosa Site (41JW8) in Jim Wells County (Jones 1986), and more recently cheno-am seeds were noted in samples from 41MV164 in Maverick County near Eagle Pass (Dering 2006).

Prickly pear. A carbonized prickly pear seed was recovered from Feature 156. In southern Texas the most common species of prickly pear is *Opuntia engelmannii*, which can grow into dense stands covering large tracts of land, although the largest stands are inland. On the coast, prickly pear can grow dense enough to provide shelter for wildlife. Prickly pear seeds have been recovered from 41UV68 in Uvalde County and 41MV164 in Maverick County (Dering 2006a and b). Cabeza de Vaca, the first European to cross Texas in the 16th century, observed that the Native Americans would squeeze the juice out of the fruit and dry the pulp and skin (Covey 1983). The stems or pads were also edible, and could be gathered and baked in an earth oven.

Indeterminate. A single carbonized plant fragment from Feature 102 has the appearance of a small maize cupule fragment. It measures about 3-mm by 1-mm and has a corky tissue adhering to one side. I can not confirm its identity, however, because it is small and fragmentary and because other parts of plants, either vegetative or reproductive, can resemble small eroded cupules. The location of this site, potentially within the Gilmore corridor, makes it necessary to mention this fragment. But I doubt that we can confirm its identity, and a single maize cupule can not demonstrate the presence of the Gilmore corridor.

Conclusion

The ephemeral nature of the occupations, the clay dunes and sand sheets, and the alkaline nature of the soils do not provide a good preservation environment for carbonized plant remains. Most of the 66 light fractions submitted for analysis were unusually small, and disturbance indicators, especially roots, were abundant. Of the 66 samples, 30 contain no carbonized plant material or only unidentifiable charcoal flecks. Three samples contain seeds, and one sample contains a fragment that resembles a maize cupule. However, small vegetative parts of plants can sometimes resemble maize cupules.

The wood charcoal is represented by five taxa, mesquite (*Prosopis glandulosa*), woody legume (*Prosopis* or *Acacia* spp.), lotebush (*Ziziphus obtusifolia*), oak (*Quercus* sp.), and brasil (*Condalia* sp.). This is the first report of brasil and lotebush wood from a southern Texas site. Two seed taxa, prickly pear and cheno-am also were identified in the samples, as noted in Table C-2. They provide the first direct archeological evidence for the use of cheno-ams and prickly pear on the middle Texas coast. The data, while sparse and not easily recovered, provide the first direct glimpse into plant use on the middle Texas coast. The presence of cheno-am seeds (fruits) is particularly interesting given its long history of use in both Mesoamerica and North America, and it adds another dimension to our understanding of land use in the Coastal Bend.

The taxa represented in the features are indicative of a local vegetation dominated by elements of the Tamaulipan Thornscrub. These elements include mesquite, several species of acacia, lotebush, brasil, prickly pear on the clay dunes, and oak on the sandy soils. Some other woody species, such as the Texas sabal palm, may have grown nearby, but I do not have direct evidence for its presence. It is likely that the area around Cayo del Oso has had similar taxa growing on it for the past 3,000 years.

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APPENDIX D:
Diatom Paleoenvironmental Analysis of Core Sediments
from the Cayo Del Oso Site, 41NU2, Nueces County, Texas

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Introduction

Diatoms are single celled algae with a siliceous cell wall. They grow in aerophilous settings, such as damp soils, wet plants and rocks, marshes, wetlands and mudflats as well as in all types of standing and flowing aquatic habitats. Their silica cells are often preserved in sedimentary deposits if water chemistry allows. Because individual taxa have specific requirements and preferences with respect to water chemistry, hydrologic conditions and substrate characteristics, diatoms are used as proxy indicators to reconstruct the paleoenvironmental and climatic changes throughout the Holocene. Fossil diatoms in lake sediments can provide sensitive high-resolution records of salinity and lake level changes, which are mainly associated with fluctuations in climate conditions such as precipitation and evaporation rates. These data are particularly useful when coupled with input about local geology and hydrology, soil characteristics, pollen, and phytoliths.

This study is part of an interdisciplinary investigation of the Cayo Del Oso Site, 41NU2. The site, a prehistoric campsite and cemetery, is located along the northwestern shore of Oso Bay from the area north of the mouth of an unnamed drainage to the shore of Corpus Christi Bay immediately west of Oso Bay. Surface and near-surface geology in the vicinity of the site consists of Holocene clay dune deposits, the Pleistocene Beaumont Clay Formation, and ephemeral ponds that developed between dune deposits and the Beaumont ridge.

The goal of this investigation is to provide information that contributes to reconstruction of the paleo-climate profile and environmental setting at the time of occupation. In the present study a total of 51 sediment samples from three cores collected at 41NU2 were submitted to Winsborough Consulting for diatom analysis. This report details the results of these analyses.

Methods

Samples were cleaned of organic material and acid-soluble minerals in preparation for microscopic analysis by boiling first in hydrogen peroxide and then in hydrochloric acid. The oxidized, decalcified material was rinsed repeatedly until a pH of about 6-7 was reached. A few drops of the cleaned material was air-dried on to glass coverslips and mounted on glass slides using NAPHRAX a resin with a high index of refraction. A slide of each sample was scanned at 1500x, until 500 diatoms were recorded, or the entire slide examined.

Results

The samples for diatom analysis were collected from three cores, numbers 2, 4, and 6, at 41NU2. There were clear differences among the kinds and numbers of diatoms found in the three different cores from the Oso site. Of the three cores, Core 2 was the most diatomaceous. Smaller numbers of diatoms were found in the other cores.

Core 2

Fifteen samples for diatom analysis were submitted from three separate segments of Core 2. All of the samples were diatomaceous and 500 diatoms were counted from all except the lowest sample, which had fewer diatoms per amount of sediment than the other samples. Table D-1 contains a complete list of samples analyzed from Core 2 with the abundances of each of the diatoms found. A total of 7,195 diatoms representing 45 taxa were recorded in the 15 samples analyzed. In decreasing order of abundance, the most commonly occurring species are *Hantzschia abundans* Lange-Bertalot, *Luticola mutica* (Kützing) Mann, *Aulacoseira italica* (Ehrenberg) Simonsen (status *semilaevis*), *Diademsia confervacea* Kützing, *Eunotia minor* (Kützing) Grunow, *Gomphonema angustatum* (Kützing) Rabenhorst, *Craticula cuspidata* (Kützing) Mann, *Luticola goeppertiana* (Bleisch) Mann, *Pinnularia borealis* Ehrenberg and *Eunotia bilunaris* (Ehrenberg) Mills.

Of these diatom species, *H. abundans*, *L. mutica*, *D. confervacea*, *E. minor*, *C. cuspidata*, *L. goeppertiana*, *P. borealis* and *E. bilunaris* are aerophils, found characteristically living as aerophils on mud and moist vegetation. They can survive under aquatic conditions but grow best when not permanently submerged. The other two species, *A. italica* (status *semilaevis*) and *G. angustatum* are aquatic, benthic species living attached to sand grains or submerged vegetation. The *semilaevis* form of *A. italica* is thought to be a resistant form adapted to the osmotic stress of drying conditions. Overall the diatoms indicate a shallow, somewhat polluted, freshwater habitat, with small to moderate amounts of salt. The pH was around 7 or above. All of the diatoms prefer circumneutral to alkaline, HARD water. The substrate was muddy, with some submerged vegetation, and the water was seasonal, decreasing substantially to moist or dry conditions for part of the year. The diatoms were heavily diluted with muddy, and silty sediments. Some of the large diatoms were broken, suggesting grazing activity by invertebrates or fish. This core represents a pond/mudflat or grassy marsh habitat that held water for part of the year and was damp to dry the rest of the time. The water temperature probably varied widely, depending on season and water depth and the diatoms are

Table D-1. Absolute Abundance of Diatoms in Core 2 from Site 41NU2

Segment	1	2	2	2	2	2	2	2	2	2	2	3	3	3	3
Zone	8	8	8	8	9	9	10	10	10	10	10	10	10	11	11
Depth (cm)	94-96	6-8	14-16	24-26	34-36	44-46	52-54	62-64	72-74	82-84	94-96	24-26	32-34	42-44	52-54
Sample #	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87
<i>Amphora veneta</i>		1													
<i>Anomoeoneis sphaerophora</i>								1					1		
<i>Aulacoseira italica</i>	1	2	14	49	65	100	91	58	79	65	105	65	34	13	8
<i>Cocconeis placentula</i>											1		1		
<i>Craticula cuspidata</i>	12	11	2		2	5	2	10	2	4	9	10	14	15	4
<i>Denticula kuetzingii</i>											2				2
<i>Diademsis confervacea</i>	2		10	16	47	63	59	81	56	78	82	59	75	58	37
<i>Diademsis contenta</i>											13		2		
<i>Diploneis puella</i>	2														
<i>Encyonema silesiaca</i>			2				2					4			
<i>Epithemia turgida</i>													2		
<i>Eunotia bilunaris</i>				6	5	8	17		8		3	4	19	2	
<i>Eunotia minor</i>	3		17	30	51	34	24	24	26	24	30	30	37	36	4
<i>Eunotia pectinalis</i>		5			2										
<i>Fragilaria fasciculata</i>		1								1	6				
<i>Gomphonema angustatum</i>	1	1	1	11	15	15	7	10	8	11	17	24	30	24	9
<i>Gomphonema gracile</i>										2	4				1
<i>Gomphonema intricatum</i>												2			
<i>Gomphonema parvulum</i>												12			
<i>Gomphonema pumilum</i>								2			2				
<i>Hantzschia abundans</i>	363	358	313	293	210	225	202	230	210	202	171	201	217	301	103
<i>Luticola goeppertiana</i>	7	7	3	6	5		1	13		36	1	5	6	2	
<i>Luticola mutica</i>	73	90	116	51	77	45	68	59	80	44	8	55	40	30	26
<i>Navicula oblonga</i>															
<i>Neidium iridis</i>	1														
<i>Nitzschia amphibia</i>	7	1	2	11	4	1	5		3		2	6	3	11	
<i>Nitzschia palea</i>									4						
<i>Nitzschia pseudocommunis</i>			3				12		2	10		2	4		
<i>Nitzschia vitrea</i>	1														
<i>Pinnularia acrosphaeria</i>										2					
<i>Pinnularia appendiculata</i>				2						5	2				
<i>Pinnularia borealis</i>	17	9	9	8	5	3	4	4	4	7	3	6	3	1	1
<i>Pinnularia braunii</i>								3			2				
<i>Pinnularia gibba</i>											4				
<i>Pinnularia microstauron</i>									2						
<i>Pinnularia subcapitata</i>		4		2	2		2								
<i>Pinnularia viridis</i>			1	2			1	1	3		4			1	
<i>Rhopalodia gibba</i>	3	4									2				
<i>Sellaphora laevissima</i>				5	2		2		4		1	6	4		
<i>Stauroneis anceps</i>	5														
<i>Stauroneis obtusa</i>	2	1	6	2	8					8		8	2		
<i>Stauroneis phoenicentron</i>			1			1					16		1	4	
<i>Surirella brebissonii</i>		3													
<i>Synedra ulna</i>		1		6			1	1	9	1	10	1	5	2	
<i>Tryblionella debilis</i>		1													
Total counted:	500	500	500	500	500	500	500	500	500	500	500	500	500	500	195

taxa adapted to these fluctuating conditions. The remaining diatom species, present in smaller numbers, confirm this paleoenvironmental interpretation.

Core 6

The 16 samples, from the 3 segments from Core 6, contained 431 diatoms representing 33 species (Table D-2). Most of these came from zone 9 of segment 1 (357 diatoms). The most abundant, in decreasing order, are *Achnantheidium*

minutissimum (Kützing) Czarnecki, *Cymbella delicatula* Kützing, *Synedra ulna* (Nitzsch) Ehrenberg, *Achnantheidium biasolettianum* (Grunow) Lange-Bertalot, *Cocconeis placentula* Ehrenberg, *Encyonopsis evergladianum* Krammer, *Denticula kuetzingii* Grunow and *Gomphonema angustatum*. These are small, aquatic, benthic, attached species known to be primary colonizers in seasonally wet, well-aerated, sunny habitats. There are almost no aerophilic or mud species, in contrast to Core 2. The diatoms indicate that the substrate was sandy, and not muddy. There was probably limited vegetation and the water was circumneutral to slightly

Table D-2. Absolute Abundance of Diatoms in Core 6 from Site 41NU2

Segment	1	1	1	1	1	1	1	2	2	2	2	3	3	3	3	3
Zone	4	6	7	8	9	9	9	10	10	11	12	13	14	15	16	17
Depth (cm)	31-33	38-39	48-49	55-56	72-73	82-83	92-93	42-43	60-62	81-82	93-94	3-4	20-21	40-41	63-64	75-78
Sample #	309	310	311	312	313	314	315	316	317	318	319	320	321	322	323	324
<i>Achnantheidium biasolettianum</i>				8	8	12										
<i>Achnantheidium minutissimum</i>					42	61	34	8			2	4	4		6	
<i>Amphora pediculus</i>					2											
<i>Cocconeis placentula</i>					9	2	2	2					12			
<i>Craticula halophila</i>						1										
<i>Cymbella affinis</i>				2												
<i>Cymbella delicatula</i>		4		12	30	2						2	2			
<i>Cymbella kolbei</i>								4			2				2	
<i>Cymbella laevis</i>					6											
<i>Denticula kuetzingii</i>					15											
<i>Diploneis puella</i>			1													
<i>Encyonema silesiacum</i>					2											
<i>Encyonopsis evergladianum</i>			2		9	2	2									
<i>Encyonopsis microcephala</i>					2							2				
<i>Eunotia pectinalis</i>					1											
<i>Fragilaria capucina</i>			3			8										
<i>Gomphonema affine</i>					2											
<i>Gomphonema angustatum</i>		6			5			1					2			
<i>Gomphonema parvulum</i>								2								
<i>Gomphonema pumilum</i>					4											
<i>Luticola mutica</i>								1								
<i>Navicula antonii</i>								4								
<i>Navicula capitatoradiata</i>					1											
<i>Navicula cincta</i>					1											
<i>Navicula cryptotenella</i>					1											
<i>Navicula menisculus</i>					2											
<i>Nitzschia amphibia</i>					9											
<i>Nitzschia amphibioides</i>					2											
<i>Nitzschia frustulum</i>								12								
<i>Rhoicosphenia abbreviata</i>		4														
<i>Sellaphora laevis</i>					1											
<i>Sellaphora stroemii</i>					2											
<i>Synedra ulna</i>				3	17		13									
Total:	0	14	6	25	173	88	51	34	0	0	4	8	20	0	8	0

alkaline (depending on amount of submerged vegetation), with rather low nutrients, but elevated salt concentrations. Evidence of grazing was not apparent.

Core 4

There were very few diatoms in Core 4 (Table D-3). Seven species were found in the 20 samples analyzed, but only 14 individual diatoms were recorded. Two of the diatoms, *Amphora pediculus* (Kützing) Grunow and *Brachysira vitrea* (Grunow) Ross are small, halophilic (salt loving), aquatic species not recorded from the other cores, but their presence is not incongruent with that of the other diatoms from the site. The assemblage as a whole suggests that the water source was fresh in origin, with elevated salt concentrations, as might be found around the saline margins of a drying habitat. The diatoms were not broken or corroded indicating that they did not travel far, but they may have been transported by winds from a nearby source.

was not one with regular, seasonally predictable rainfall. Rain was episodic and only those diatoms capable of lasting through dry intervals and growing rapidly during wet periods were able to persist in this somewhat osmotically stressed habitat. There was drinkable water available seasonally during the interval identified as the pond deposit. It was hard water but there is no evidence of pollution or brackish water conditions. The absence of planktonic diatoms means that the pond was shallow, reasonably clear and probably somewhat low in nutrients. The few diatoms found in the sandy material of Cores 4 and 6 indicate standing water that could have been associated with a clay-lined depression or swale in the sands, in contrast to the muddy pond sediments. The water source supplying the diatoms found in Cores 4 and 6 may not have lasted more than a few weeks, however, and there was no soil or mud to support any aerophilic taxa. The greater diversity in diatom species in the pond deposit is probably due to the presence of more different microhabitats, greater depth, and longer residence time of the water, providing more available niches for diatoms to colonize, within the broader pond environment.

Conclusions

In summary, the diatoms associated with the pond interval of Core 2 at the Oso site indicate that the general climate

Table D-3. Absolute Abundance of Diatoms in Core 4 from Site 41NU2

Segment	1	1	1	1	1	1	1	1	2	2	2	2	2	2	2	2	3	3	3	
Zone	5	6	7	8	9	10	12	13	15	16	17	18	19	23	23	24	25	26	27	28
Depth (cm)	39-44	44-49	52-54	54-60	65-70	74-76	83-87	90-94	6-10	10-13	23-30	34-38	42-46	63-70	78-88	92-96	96-100	20-30	30-40	50-60
Sample #	351	352	353	354	355	356	357	358	359	360	361	362	363	364	365	366	367	368	369	370
<i>Achnanthydium biasolettianum</i>																				3
<i>Achnanthydium minutissimum</i>																2				
<i>Amphora pediculus</i>															2					
<i>Brachysira vitrea</i>		2																		
<i>Fragilaria fasciculata</i>		1																		
<i>Hantzschia abundans</i>	1									1										
<i>Synedra ulna</i>		1																		1
Total counted:	1	4	0	0	0	0	0	0	0	1	0	0	0	0	2	2	0	0	0	4

APPENDIX E:
**Late Holocene Paleoenvironmental History of Cayo Del Oso Bay,
Nueces County, Texas: Geoarchaeological Implications**

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Introduction

Understanding human occupation in southern Texas requires an understanding of the current and past conditions of the area. Environmental factors do not necessarily rule the course of human adaptation, but they do impose constraints and set parameters that affect people's lives and their use of resources. This paper focuses on the micropaleontological analysis of three cores collected within a small water body at Cayo del Oso Bay, Nueces County, Texas. To achieve this goal Terra Nostra Earth Sciences Research staff analyzed the microfossil contents of three cores (2, 4, and 6) obtained at site 41NU2 in the summer of 2007. The results were compared and contrasted with previous work in the region to establish the paleoenvironmental patterns that characterized the area during the late Holocene. The sections below summarize the main features of the three biological groups found during this research: ostracodes, foraminifera, and molluscs.

Marginal marine environments (estuaries and lagoons) provide a unique source of information about short- (e.g., seasonality, human activity) and long-term (e.g., isostatic or eustatic sea level change) continental and marine cycles that have affected human adaptation. Sedimentation in estuaries and lagoons is influenced by a complex combination of tidal currents, oceanic waves, locally generated waves, river discharge (into estuaries), precipitation, temperature, and local flora and fauna (Clifton 1982). Also important are geologic agents like tectonism, subsidence, isostatic and eustatic changes of sea level, climate change, and human activities that have altered the hydrology, drainage, and topography.

Lagoons differentiate from estuaries both physically and ecologically. Lagoons are closed marine shallow embayments, less affected by fluvial discharge than estuaries (although, many estuaries form lagoonal environments). A major difference between the two is that whereas estuaries are elongated perpendicular to the coastline and form deltas, lagoons are elongated parallel to the coastline separated from the ocean by sand barriers. Ecologically, lagoons include a mostly marine, poorly diverse faunal assemblage consisting of ostracodes, foraminifera, diatoms, molluscs, and other groups. Estuaries, by contrast, contain a poorly abundant but diverse assemblage that include both marine and nonmarine species of many groups that have representatives in both realms, mainly ostracodes, molluscs, and diatoms.

Ostracodes Biology and Ecology

The Ostracoda belong to the phylum Arthropoda (Figure E-1). Microcrustaceans provided with a calcium carbonate carapace, ostracodes occur in almost any water environment

150 x

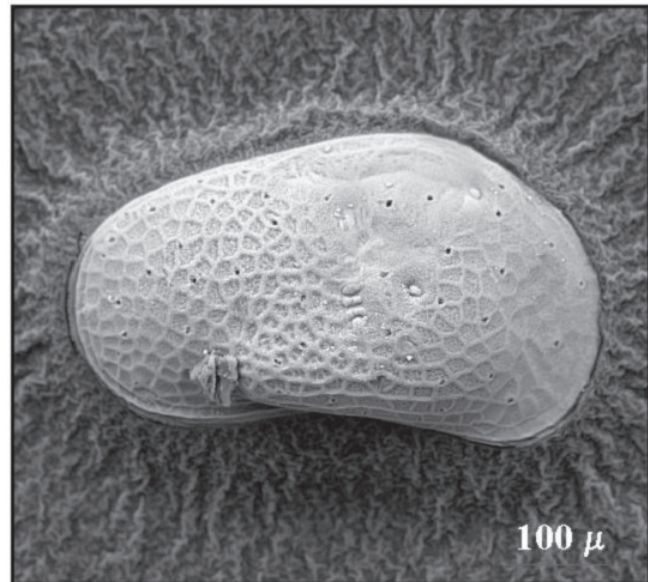


Figure E-1. *Limnocythere staplini*, euryhaline continental water ostracode common in marginal marine systems. Female, right valve view. (Image by Manuel R. Palacios-Fest, TNESR.)

from the ocean to rivers, lakes, ponds, springs and even in some tropical forest soils like in the Amazon. As other crustaceans, ostracodes require to grow by shedding their exoskeleton about nine times before becoming adults. During the interval before calcification they grow to reach the next stage. Both adult and juvenile valves (named instars) may be preserved as fossils, and both are useful for paleontologic studies. The carapace morphology and shape are both environmentally and genetically controlled (Pokorný 1978).

Ostracodes are widely distributed in the marine and freshwater environments, normally under well-oxygenated conditions. Marginal marine species are also abundant. Although all modern families have marine and freshwater representatives, distinctive species are characteristic of each environment. Only brackish-water ostracodes occur in continental saline or marginal marine environments. Water chemistry and the sediment and nutrient properties of the environment may impose a mechanism of control on them. Estuarine ostracodes may be represented by a combination of marine, brackish- and freshwater species under certain circumstances.

Marinov (1964; *in*: Puri *et al.* 1969) recognized four categories of ostracode distribution according to their origin in the Black Sea estuaries, as follows: 1) Freshwater species which have also been found in the brackish-water basins. Most of the species found belonged to the genera *Ilyocypris*, *Candona*, *Limnocythere*, *Cypridopsis* and *Darwinula*; 2)

typical brackish-water species, mainly represented by several species of the genus *Cyprideis* and some other genera like *Hemicythere* and *Potamocypris*; 3) ultrahaline species represented by the genus *Eucypris*; and 4) Sea-euryhaline species which are open sea forms that have adapted themselves in mesohaline waters represented among other genera by *Cytheridea*, *Pterygocythereis*, *Hemicytherura*, *Loxoconcha*, *Xestoleberis* and *Paradoxostoma*. A similar pattern from Texas is discussed in this study, although the freshwater assemblage is absent and the ultrahaline is represented by a single specimen of the genus *Limnocythere*.

Ostracode studies of southern Texas are numerous, among which of particular interest to this study are the works of Van den Bold (1977), Cronin (1986), Garbett and Maddocks (1979), King and Kornicker (1970), and Swain (1955). For example, Cronin (1986), studying the micropaleontological record of six cores from Mustang Island, Corpus Christi Bay, identified ostracode assemblages that permitted him to infer episodes of freshwater-fluvial input alternating with upper bay, mid-bay, lower bay-inlet, and sublittoral environments during the Quaternary. His work is of significant value to this study because it is the first attempt to reconstruct the Quaternary paleoenvironments of deposition in Corpus Christi Bay.

Foraminifera Biology and Ecology

The Foraminifera belong to the phylum Sarcomastigophora (Figure E-2). Foraminifera are found in all marine environments, they may be planktic or benthic in mode of life. These unicellular organisms are provided with a calcareous or agglutinated test that preserves to the geologic record. The morphology of foraminifera tests varies enormously, but in terms of classification two features are important: chamber arrangement and aperture type (for details see Boersma 1978; <http://www.ucl.ac.uk/GeolSci/micropal/foram.html> 2007). The benthic forms are the most common and abundant in marginal marine environments (e.g., estuaries, lagoons).

Benthic foraminifera are most commonly used for paleobathymetric reconstructions because of the specific depth range of many species. Some of which are restricted to nearshore, reefal, or lagoonal environments; while other are limited to deeper open ocean conditions. These organisms are divided into morphogroups based on the test shape allowing researchers to identify paleohabitats and substrates. For example, infaunal species tend to be elongate and streamlined in order to burrow into the sediments; whereas epifaunal species are more globular with one relatively flatter side in order to facilitate movement on top of the substrate. In addition, the test wall type (e.g., porcelaneous, hyaline, agglutinated) also indicates paleodepths and salinity (Boersma 1978; <http://www.ucl.ac.uk/GeolSci/micropal/foram.html> 2007).



Figure E-2. *Ammonia beccarii*, benthic foraminifera common in nearshore, lagoonal, and estuarine environments. Umbilical view. (SEM image courtesy of Elisabeth Alve, University of Oslo, originally published in the *Journal of Foraminiferal Research* 31:1, used with permission.)

In a recent study, Ishman (1997) reported two main biofacies in the Biscayne Bay, Florida. An open-bay coarse-sediment, grassy biofacies containing abundant miliolid taxa and a restricted biofacies characterized by abundant *Ammonia parkinsoniana*, *Elphidium galvestonense*, and species associated with high salinity tolerance. A similar pattern from Texas is discussed in this study.

As for ostracodes, the work on foraminifers is abundant along the Texas Gulf Coast. The studies of Phleger (1965, 1970) are key for the paleoenvironmental reconstruction of Cayo del Oso Bay because they provide the modern distribution patterns of marginal marine foraminifers in the nearby Galveston Bay, an area dominated by *Ammonia beccarii*.

Mollusc Biology and Ecology

Molluscs belong to the phylum Mollusca (Figure E-3). It is a very large and diverse group of soft body organisms that include bivalves and gastropods, as well as other forms characterized by an internal shell. The first two are of primary interest to paleoenvironmental studies because they preserve well into the geologic record; therefore, this summary will concentrate in these groups. Molluscs are triploblastic protostomes (incipient bilateral symmetry with three germ layers: ectoderm, mesoderm, and endoderm) and many demonstrate bilateral symmetry. They have a mantle lining



Figure E-3. *Physa virgata*, small freshwater gastropod common in continental and marginal marine environments. (Image by Manuel R. Palacios-Fest, TNESR.)

the shell and a muscular foot that in most species is used for locomotion. In most molluscs the mantle secretes a calcium carbonate shell that preserves well into the geologic record since the Cambrian (Barnes 1987; Brusca and Brusca 1990).

Based on their diversity, molluscs are adapted to a wide variety of environments, both aquatic and terrestrial. In marginal marine environments both groups are equally important because their occurrence may indicate transgression/regression events over time. Furthermore, aquatic forms may be represented by marine and nonmarine organisms that are also important indicators of environmental change. Sensitive to pH, alkalinity, salinity, and temperature, molluscs are restricted by the conditions prevailing in a given area. Although modern families have marine, freshwater, and terrestrial representatives, distinctive species characterize each environment. Water chemistry and the sediment and nutrient properties of the environment may impose a mechanism of control on them. Terrestrial species may tolerate water-saturated soils; other may thrive in drier conditions (Barnes 1987; Brusca and Brusca 1990).

Purpose of Study

This study documents the occurrence of lagoonal or estuarine biofacies in site 41NU2 at Cayo del Oso Bay, Nueces County, Texas. Marine and continental assemblages randomly alternated in the stratigraphic record of Cayo del Oso during the last ~3000 years, before that, marine species dominated the record. Ostracode biofacies analysis defined

environmental change during the late Holocene. Two main questions addressed in this research, are: 1) What was the aquatic environment met by prehistoric people around Cayo del Oso? And 2) Did the lagoon receive significant freshwater input over time?

Area of Study

Site 41NU2, Cayo del Oso, Nueces County, Texas, lies on clay dunes along the northwestern shore of Oso Bay and extends from a small creek that empties into the western side of Oso Bay approximately 920 m north to Corpus Christi Bay (see Figure 1-2, this volume). The site's width is known only to the northern border, on a bluff overlooking Corpus Christi Bay. At that point the site is approximately 200 m wide (Tomka et al. 2007).

Geologically, Corpus Christi Bay and the Oso Bay area began forming about 5000 years B.P. as the Padre Island sand barriers coalesced, a process that was completed about 2500 years B.P. forming Laguna Madre (Fisk 1959:123; Morton and Paine 1984:12). A shallow (<4 m deep) bay formed (Harwood 1980) receiving freshwater discharge from the Nueces River and Oso Creek, giving rise to the estuarine conditions prevailing today (Ricklis 1995:142). In historic time, two channels (Aransas Pass and Fish Pass) were opened to the ocean connecting Corpus Christi Bay to the Gulf of Mexico (Morton and Pieper 1977).

At the southwestern end of Corpus Christi Bay is Oso Bay, an enclosed, soft-sediment, estuary where the Oso Creek discharges (Tomka et al. 2007:9). The bay is less than 1 m deep, connected to Corpus Christi Bay by a small pass east of Ward Island. A small wind-flat, covering approximately 28 hectares, lies west of Ward Island, called False Oso or Blind Oso Bay. Site 41NU2 is located on the western shore of the False Oso Bay (Tomka et al. 2007:9).

Southern Texas is characterized by a mean annual precipitation of 28.86 in (73 cm) distributed in two seasons: late spring/early summer, and fall (Kier and White 1978). Evaporation exceeds precipitation with rates between 90 and 115 cm/year (occasionally exceeding 150 cm/year). The atmospheric summer temperatures average 33.3°C, while winter lows average 8.3°C, while water temperatures in Corpus Christi Bay range from 7°C in the winter to 29°C in the summer (Harwood 1980).

Corpus Christi Bay is a hypersaline to brackish-water system where salinity ranges from 11 parts per thousand (ppt) to 45 ppt, with a mean value around 30 ppt affected by the southeasterly

winds during the summer and strong northerly fronts during the winter (Chabreck 1990). This condition has a great influence on the bay's ecology (for details see Tomka et al. 2007).

Three of six cores collected were used for this study (2, 4, and 6) (see Figure 5-1, 5-2 and 5-4, this volume). Core 2 was collected at the northwest end of the area of study in a freshwater pond near site 41NU2. The pond formed on and at the foot of the upland Beaumont Formation and behind the clay dunes well above sea level (Frederick, written communication, January 25, 2008). It possibly received some eolian component from Cayo del Oso Bay, runoff from the Beaumont Creek, and overwash from the bay and Gulf of Mexico during hurricanes. Cores 4 and 6 were collected along the north and south dunes that developed against a remnant ridge of Pleistocene Beaumont Clay exposed to the surface between them (Meissner, written communication, January 23, 2008).

Materials and Methods

The three cores (2, 4, and 6) were subsampled and the samples shipped to Terra Nostra Earth Sciences Research, LLC in Tucson, Arizona for micropaleontological analysis. A total of 45 samples ranging in size from 16 to 56 g were processed for this analysis.

All samples were processed for micro-invertebrate recovery following Palacios-Fest's (1994). Samples were weighed using an electronic balance and disaggregated in hot distilled water with Alconox (Sodium Hexametaphosphate). The sediments were then left to cool and let to sit at room temperature for eight days. They were washed using a set of three sieves (>1 mm, >106 μm , and >63 μm) for separating the three sand fractions, and rinsed with distilled water. The residues were let dry at room temperature for 48 hours and weighed to estimate the content of micro-invertebrates per gram of sediment. The rough particle-size analysis obtained during this phase of the study permits making an estimate of sediment input over time at Cayo del Oso Bay. Residues were examined under a low power stereoscopic microscope. Routine micropaleontologic analysis was performed to determine fossil content and faunal composition (ostracodes, foraminifera, and molluscs). The focus of the present study is ostracodes. Foraminifera and molluscs provided complementary data to support the ostracode records. Ostracode abundance ranged from extremely rare to moderately abundant¹.

Ostracode paleoecologic assemblages were defined and grouped according to species abundance following Palacios-

Fest et al. (2006), Benson and Coleman (1963), and McKenzie and Swain (1967). To reconstruct the late Holocene history of Cayo del Oso Bay, a qualitative estimate of environmental change was obtained assigning to each species a relative value as determined by equation (1):

$$\text{SI} = [3(\% \textit{Cyprideis salebrosa}) + 2(\% \textit{Cyprinotus glaucus}) + (\% \textit{Candona rawsoni})] - [4(\% \textit{Aurila laevicula}) + 3(\% \textit{Radimella wantlandi}) + 2(\% \textit{Perissocytheridea rugata}) + (\% \textit{Haplocytheridea proboscidiata}) - (\% \textit{Loxococoncha avellana}) + 2(\% \textit{Cytherura sandbergi}) + 3(\% \textit{Xestoleberis banda})]$$

where the marine species (see Table E-4) were grouped from sublittoral-tolerant to upper bay-tolerant species and subtracted from the limnic species. The limnic species were ordered from higher salinity-tolerant to lower salinity-tolerant (see Table E-3).

In addition to the faunal composition, the taphonomic features were used to recognize degrees of transport and burial characteristics, like desiccation and sediment compaction. The rates of fragmentation, abrasion, and disarticulation are indicators of transport. Commonly, these parameters show increasing damage with increasing transport. One must be cautious in using these criteria; however, the nature of the deposits suggests that micro-invertebrates reflect the ecological properties of Cayo del Oso Bay. Deposition characteristics were determined by encrustation, coating, the redox index which reflect authigenic mineralization, wave action, and burial conditions, respectively. Disarticulation (C/V) and adulthood (A/J) ratios were used as indicators of biocenosis (Palacios-Fest et al. 2001).

The Sedimentary and Faunal Records

The cores used for the micropaleontological analysis are described by Frederick (Appendix F, this volume) the reader is invited to review that chapter for further reference. Table E-1 and Figure E-4 show the particle-size analysis obtained during the sample washing process. Core 2 consisted of very dark gray (5Y 3/1) to gray (5Y 6/1) clay (>90%) except for a few intervals where silt increased (>84%<90%). Core 4 had a more variable particle-size ranging from gray (5Y 6/1) to white (5Y 8/1) but mostly light gray (5Y 7/1) sandy silt to clay. Core 6 consisted of light gray (5Y 7/1) occasionally alternating with gray (5Y 6/1) or white (5Y 8/1) sandy silt to clay.

¹ Abundance explanation (number of individuals): Extremely abundant (>1001), very abundant (>501<1000), abundant (>101<500), moderately abundant (>51<100), common (>21<50), rare (>6<20), and extremely rare (<5).

Table E-1. Sample Identification, Radiocarbon Dates and Textural Classification of Sediment Residues Including Munsell's Color Code from Cores 2, 4, and 6 of Cayo del Oso Bay, Nueces County, Texas

Core #	Segment #	Zone #	Sample ID	Depth (cm)	Total Depth (cm bgs)	Conventional 14C Age* (Years B.P.)	Sedimentation Rates (cm yr-1)	Bulk Wt. (g)	Fraction Wt. (g)	>1mm (g)	>106µm (g)	>63µm (g)	<63 µm (g)	>1mm (%)	>106µm (%)	>63µm (%)	<63 µm (%)	Textural Classification	Munsell's Color				
2	1	8	118	93-96	95			32.9	5.1	0.1	2.6	2.4	27.8	0.3	7.9	7.3	84.5	Clay	Gray	5Y 6/1			
			67	96-98	97	550±40	0.05																
		2		119	4-8	112			36.7	4.5	0.2	1.5	2.8	32.2	0.5	4.1	7.6	87.7	Clay	Gray	5Y 6/1		
				120	14-17	120			28.8	2.0	0.0	0.6	1.4	26.8	0.0	2.1	4.9	93.1	Clay	Gray	5Y 6/1		
				121	23-26	129			33.8	1.2	0.0	0.8	0.4	32.6	0.0	2.4	1.2	96.4	Clay	Gray	5Y 6/1		
			9	122	32-35	138			36.6	2.5	0.0	0.8	1.7	34.1	0.0	2.2	4.6	93.2	Clay	Very dark gray	5Y 3/1		
				69	36-38	140	1300±40	0.02															
				123	42-44	145			37.8	3.1	0.1	1.2	1.8	34.7	0.3	3.2	4.8	91.8	Clay	Very dark gray	5Y 3/1		
				124	52-56	155			31.9	2.2	0.1	0.5	1.6	29.7	0.3	1.6	5.0	93.1	Clay	Very dark gray	5Y 3/1		
				71	28-30	160	2240±40	0.04															
			125	62-64	163			43.4	3.0	0.1	0.8	2.1	40.4	0.2	1.8	4.8	93.1	Clay	Dark gray	5Y 4/1			
			126	70-73	170			35.7	2.4	0.3	0.8	1.3	33.3	0.8	2.2	3.6	93.3	Clay	Dark gray	5Y 4/1			
			127	80-82	180			47.3	2.8	0.0	1.1	1.7	44.5	0.0	2.3	3.6	94.1	Clay	Very dark gray	5Y 3/1			
			128	90-93	190			54.7	4.9	0.2	1.2	3.5	49.8	0.4	2.2	6.4	91.0	Clay	Very dark gray	5Y 3/1			
			3	10	129	24-26	232			36.3	3.7	0.3	0.8	2.6	32.6	0.8	2.2	7.2	89.8	Clay	Dark gray	5Y 4/1	
					130	32-34	239			30.4	4.3	0.3	1.1	2.9	26.1	1.0	3.6	9.5	85.9	Clay	Dark gray	5Y 4/1	
				11		131	40-42	244			36.8	5.1	0.2	1.2	3.7	31.7	0.5	3.3	10.1	86.1	Clay	Dark gray	5Y 4/1
						132	49-51	252			51.5	7.2	0.7	2.5	4.0	44.3	1.4	4.9	7.8	86.0	Clay	Gray	5Y 5/1
						72	56-58	259	2970±40														
			4	1	6	240	48-52	50			53.5	17.3	0.2	8.4	8.7	36.2	0.4	15.7	16.3	67.7	Sandy silt	Light gray	5Y 7/1
241	60-63	60						55.7	16.3	0.1	5.5	10.7	39.4	0.2	9.9	19.2	70.7	Sandy silt	White	5Y 8/1			
	9			153	65-68	65	1060±40	0.04															
				242	70-74	70			33.1	9.2	0.3	3.2	5.7	23.9	0.9	9.7	17.2	72.2	Silty clay	White	5Y 8/1		
	11			243	80-83	80			47.1	16.7	0.6	8.6	7.5	30.4	1.3	18.3	15.9	64.5	Sandy silt	Light gray	5Y 7/1		
				244	88-90	90			47.5	14.0	0.4	6.4	7.2	33.5	0.8	13.5	15.2	70.5	Sandy silt	Light gray	5Y 7/1		
	13			245	95-97	95			40.6	14.6	1.5	7.0	6.1	26.0	3.7	17.2	15.0	64.0	Sandy silt	Light gray	5Y 7/1		

* Radiocarbon dates obtained from humate samples. For details see Tomka et al. (This volume)

Table E-1. Continued...

Core #	Segment #	Zone #	Sample ID	Depth (cm)	Total Depth (cm bgs)	Conventional 14C Age* (Years B.P.)	Sedimentation Rates (cm yr-1)	Bulk Wt. (g)	Fraction Wt. (g)	>1mm (g)	>100µm (g)	>63µm (g)	<63 µm (g)	>1mm (%)	>100µm (%)	>63µm (%)	<63 µm (%)	Textural Classification	Munsell's Color	
4	2																			
			156	13	108	2110±40	0.05													
		16	247	15-17	111			59.1	14.9	0.6	6.0	8.3	44.2	1.0	10.2	14.0	74.8	Clayey sandy	Gray	5Y 6/1
		17	248	22-24	117			38.8	6.2	0.3	2.8	3.1	32.6	0.8	7.2	8.0	84.0	Clayey silt	Light gray	5Y 7/1
		18	250	38-42	130			25.5	4.0	0.1	1.6	2.3	21.5	0.4	6.3	9.0	84.3	Clay	Light gray	5Y 7/1
		19	251	46-49	135			18.5	3.1	0.0	1.6	1.5	15.4	0.0	8.6	8.1	83.2	Clay	Light gray	5Y 7/1
		22	157	59-61	145	2880±40	0.07													
		23	252	75-78	160			34.4	7.4	0.1	3.4	3.9	27.0	0.3	9.9	11.3	78.5	Clayey silt	Light gray	5Y 7/1
			253	85-88	170			35.9	7.1	0.1	3.2	3.8	28.8	0.3	8.9	10.6	80.2	Clay	Light gray	5Y 7/1
		25	254	98-	180			42.1	8.8	0.5	4.3	4.0	33.3	1.2	10.2	9.5	79.1	Clayey silt	Light gray	5Y 7/1
	3	26	158	14-16	195	3620±40														
		26	255	17-20	200			16.1	6.4	0.0	2.6	3.8	9.7	0.0	16.1	23.6	60.2	Sandy silt	Light gray	5Y 7/1
		27	256	40-43	240			28.3	10.4	0.1	4.4	5.9	17.9	0.4	15.5	20.8	63.3	Sandy silt	Light gray	5Y 7/1
		28	257	55-58	255			33.5	11.9	0.4	4.3	7.2	21.6	1.2	12.8	21.5	64.5	Sandy silt	White	5Y 8/1
6	1	5	295	34-36	35			22.5	5.7	0.1	2.8	2.8	16.8	0.4	12.4	12.4	74.7	Clayey sandy	Light gray	5Y 7/1
		6	296	42-45	42			19.2	7.3	0.3	2.3	4.7	11.9	1.6	12.0	24.5	62.0	Sandy silt	White	5Y 8/1
		7	259	44-45	45	1280±40	0.18													
		8	297	50-53	50			28.4	5.2	0.3	1.4	3.5	23.2	1.1	4.9	12.3	81.7	Clay	Gray	5Y 6/1
		9	298	67-70	70			34.7	5.5	0.1	2.3	3.1	29.2	0.3	6.6	8.9	84.1	Clay	Light gray	5Y 7/1
			299	77-80	80			29.0	4.0	0.1	1.8	2.1	25.0	0.3	6.2	7.2	86.2	Clay	Light gray	5Y 7/1
			300	87-90	90			32.5	5.1	0.1	2.3	2.7	27.4	0.3	7.1	8.3	84.3	Clay	Light gray	5Y 7/1
	2	10	260	42-43	133	1770±40	0.04													
			301	56-58	148			18.9	5.1	0.1	2.2	2.8	13.8	0.5	11.6	14.8	73.0	Clayey sandy	Light gray	5Y 7/1
		11	302	78-80	170			27.2	7.9	0.2	4.5	3.2	19.3	0.7	16.5	11.8	71.0	Sandy silt	Light gray	5Y 7/1
		12	261	86-87	177	2760±40	0.07													
			303	91-93	183			26.9	8.4	0.4	3.3	4.7	18.5	1.5	12.3	17.5	68.8	Sandy silt	Light gray	5Y 7/1
	3	13	304	9-13	203			28.9	7.1	0.1	2.0	5.0	21.8	0.3	6.9	17.3	75.4	Clayey sandy	Pale Yellow	2.5Y
		14	305	22-25	228			29.8	7.9	0.2	3.3	4.4	21.9	0.7	11.1	14.8	73.5	Clayey silt	Light gray	2.5Y
		15	262	40-42	245	3770±40														
			306	45-48	251			35.8	8.1	0.2	2.8	5.1	27.7	0.6	7.8	14.2	77.4	Sandy silt	Light gray	5Y 7/1
			307	55-58	261			32.9	8.4	0.3	3.1	5.0	24.5	0.9	9.4	15.2	74.5	Sandy silt	Light gray	2.5Y
		17	308	73-75	278			32.6	7.7	0.4	2.8	4.5	24.9	1.2	8.6	13.8	76.4	Sandy silt	White	2.5Y

* Radiocarbon dates obtained from humate samples. For details see Tomka et al. (This volume)

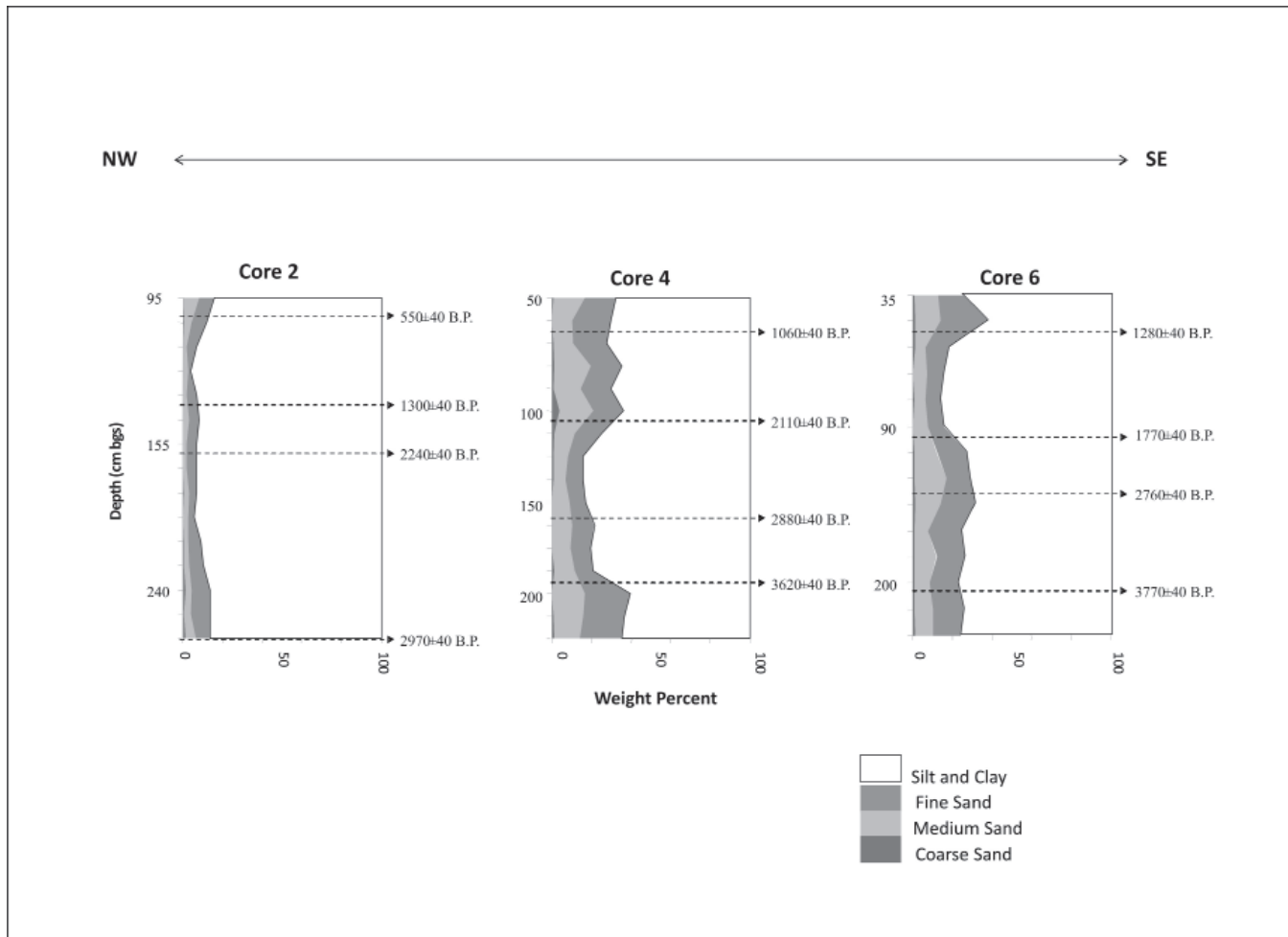


Figure E-4. Particle-size analysis from ostracode samples (cores 2, 4, and 6) for establishing the the site paleohydraulics with the micro-invertebrate and foraminifera populations.

A total of 45 sediment samples were analyzed for micro-invertebrates of which 34 contained microfossils. Limnic ostracodes occurred in three samples of core 2 (15 samples), 11 of core 4 (16 samples), and eight of core 6 (14 samples). Marine ostracodes did not occur in core 2, but were present in 11 samples of core 4 and 12 of core 6. Foraminifera did not occur in core 2, except for three specimens at sample 129, but were present in all samples of cores 4 and 6. Molluscs occurred only once in core 2 but appeared in 13 samples of core 4 and ten samples of core 6.

Table E-2 shows the faunal composition and taphonomic features (fragmentation, abrasion, encrustation, coating, redox index, and shell (or test) color) recorded at the three cores. Ostracodes, foraminifera, and molluscs occurred in Cayo del Oso Bay. Foraminifera were the most diverse and abundant biologic group in the cores (9 species, 3-606 specimens). This group occurred in cores 4 and 6; core 2 only contained nonmarine ostracodes and molluscs. Molluscs were

scarce both in diversity and abundance (1-4 species, 1-17 specimens). Ostracodes were poorly to moderately diverse and abundant (1-10 species, 1-87 specimens).

Table E-3 documents the list of ostracode species identified from Cayo del Oso cores and the ecological parameters controlling their occurrence in the basin. Thirteen species were identified, five of which are common to fresh- to saline continental waters: *Cyprideis salebrosa* (Van den Bold, 1963), *Cyprinotus glaucus* (Furtos, 1933), *Candona rawsoni* (Tressler, 1957), *Cypridopsis* sp., and *Limnocythere* sp. cf. *L. floridensis* Keyser, 1975. The remaining eight species are all marine (common in estuarine, lagoonal, and nearshore open marine environments): *Perissocytheridea rugata* Swain, 1955, *Radimella wantlandi* Teeter, 1975, *Cytherura sandbergi* Morales, *Loxoconcha avellana* (Brady), *Aurila laevicula* (Edwards, 1944), *Xestoleberis banda* (Benson, 1959), *Haplocytheridea proboscidiata* (Edwards), and *Cushmanidea* sp.

Table E-2. Micropaleontologic Composition of Cores 2, 4, and 6 from Cayo del Oso Bay, Nueces County, Texas, Including Taphonomic Parameters

Sample ID	Depth (cm)	Total Depth (cm bgs)	Conventional 14C Age* (Years B.P.)	Bulk Wt. (g)	Fraction Wt. (g)	Foraminifera	Ostracodes	Mollusks	Plant Debris	Taphonomy:					Shell Color
										Fragmentation	Abrasion	Encrustation	Coating	Redox Index	
118	93-96	95		32.9	5.1		4	2		30	15	10	0	0	White
67	96-98	97	550±40												
119	4-8	112		36.7	4.5		1			30	15	10	0	0	White
120	14-17	120		28.8	2										
121	23-26	129		33.8	1.2		1			0	0	0	0	0	Clear
122	32-35	138		36.6	2.5										
69	36-38	140	1300±40												
123	42-44	145		37.8	3.1										
124	52-56	155		31.9	2.2										
71	28-30	160	2240±40												
125	62-64	163		43.4	3										
126	70-73	170		35.7	2.4										
127	80-82	180		47.3	2.8										
128	90-93	190		54.7	4.9										
129	24-26	232		36.3	3.7	3				0	0	0	0	0	White
130	32-34	239		30.4	4.3										
131	40-42	244		36.8	5.1										
132	49-51	252		51.5	7.2										
72	56-58	259	2970±40												
240	48-52	50		53.5	17.3	606	13	15		15	10	0	5	0	Clear
241	60-63	60		55.7	16.3	473	12	2	X	10	10	0	5	0	Clear
153	65-68	65	1060±40												
242	70-74	70		33.1	9.2	110	7	2	X	15	10	0	5	0	White
243	80-83	80		47.1	16.7	95	2		X	10	5	0	5	0	Clear
244	88-90	90		47.5	14	44		3	X	15	10	0	5	0	White
245	95-97	95		40.6	14.6	14		2	X	30	10	0	0	0	White
156	13	108	2110±40												
247	15-17	111		59.1	14.9	74	2	13	X	15	10	0	5	0	White
248	22-24	117		38.8	6.2	83	16	6	X	15	10	10	5	0	White
250	38-42	130		25.5	4	143	22	5	X	10	5	0	5	0	Clear
251	46-49	135		18.5	3.1	87	13	4	X	10	5	0	5	0	Clear
157	59-61	145	2880±40												
252	75-78	160		34.4	7.4	594	87	17	X	5	5	0	5	0	Clear
253	85-88	170		35.9	7.1	110	12	15	X	5	5	0	5	0	Clear
254	98-100	180		42.1	8.8	545	67	5	X	5	5	0	5	0	Clear
158	14-16	195	3620±40												
255	17-20	200		16.1	6.4	9			X	5	0	0	0	0	Clear
256	40-43	240		28.3	10.4	13		1	X	15	10	0	5	0	Clear
257	55-58	255		33.5	11.9	3			X	15	10	0	5	0	Clear
295	34-36	35		22.5	5.7	222	1	1	X	5	0	0	0	0	Clear
296	42-45	42		19.2	7.3	77	2	1	X	5	0	0	0	0	Clear
259	44-45	45	1280±40												
297	50-53	50		28.4	5.2	7			X	0	0	0	0	0	Clear
298	67-70	70		34.7	5.5	387	1	3	X	5	5	0	0	0	Clear
299	77-80	80		29	4	244	9	6	X	5	0	0	0	0	Clear
300	87-90	90		32.5	5.1	281	7	3	X	5	5	0	0	0	Clear
260	42-43	133	1770±40												
301	56-58	148		18.9	5.1	144	6	2		5	5	0	0	0	Clear
302	78-80	170		27.2	7.9	162	11	4		5	5	0	0	0	Clear
261	86-87	177	2760±40												
303	91-93	183		26.9	8.4	149	2			5	5	0	0	0	Clear
304	9-13	203		28.9	7.1	60	1			5	5	0	0	0	Clear
305	22-25	228		29.8	7.9	55	1	1	X	5	5	0	0	0	Clear
262	40-42	245	3770±40												
306	45-48	251		35.8	8.1	47	3	1		5	5	0	0	0	White
307	55-58	261		32.9	8.4	21	3	2		5	5	0	0	0	White
308	73-75	278		32.6	7.7	17				5	5	0	0	0	White

* Radiocarbon dates obtained from humate samples. For details see Tomka et al. (This volume)

Table E-3. Ecological Requirements of Ostracode Species Recovered from Site 41NU2, Cayo del Oso Bay, Nueces County, Texas

Species	Habitat	Permanence	Temperature		Salinity*	Chemistry*	
<i>Cyprideis salebrosa</i> (Van den Bold 1963)	Springs, streams, lakes	Permanent or ephemeral	2-32°C	Eurythermic	100-4,000 mg L-1	0.10-50 meq L-1	Freshwater to Ca-rich
<i>Cyprinotus glaucus</i> (Furtos, 1933)	Streams, lakes, ponds	Permanent or ephemeral	6-18°C	Eurythermic	100-4,000 mg L-1	0.10-50.0 meq L-1	Freshwater to Ca-rich
<i>Candona rawsoni</i> (Tressler, 1957)	Lakes, ponds	Permanent	24-32°C	Eurythermic	500-5,000 mg L-1	0.10-4.0 meq L-1	Freshwater to Ca-rich
<i>Limnocythere</i> ? sp.	Springs, streams, lakes	Permanent or ephemeral	NA	Eurythermic	NA	NA	Freshwater to Ca-rich
<i>Cypridopsis</i> sp.	Springs, streams, lakes	Permanent or ephemeral	NA	Eurythermic	NA	NA	Freshwater to Ca-rich
<i>Perissocytheridea swaini</i> (Benson and Kaesler, 1963)	Estuarine or lagoonal	Permanent					
<i>Radimella confragosa</i> (Edwards)	Estuarine or lagoonal	Permanent					
<i>Cytherura sandbergi</i> Morales	Estuarine or lagoonal	Permanent					
<i>Loxoconcha avellana</i> (Brady)	Estuarine or lagoonal	Permanent					
<i>Aurila conradi</i> (Howe and McGuirt)	Estuarine or lagoonal	Permanent					
<i>Cushmanidea</i> sp.	Estuarine or lagoonal	Permanent					
<i>Xestoleberis banda</i> Benson, 1959	Estuarine or lagoonal	Permanent					
<i>Haploocytheridea probosciduala</i> (Edwards)	Estuarine or lagoonal	Permanent					

* *NANODE*

Table E-4 shows the total population, total and relative abundance, disarticulation and adulthood ratios by species, and the paleosalinity index generated by equation (1). The paleosalinity index is a dimensionless number with no true salinity value obtained to produce the salinity patterns shown in Figures E-5 through E-7.

Table E-5 displays the list of foraminifera species identified from cores 4 and 6: *Ammonia beccarii* (Linné), *Elphidium galvestonense* Kornfeld, *Elphidium gunteri* Cole, 1931, *Triloculina bicarinata* (d'Orbigny), *Triloculina linneiana* (Brady, 1884), *Quinqueloculina laevigata* (d'Orbigny), *Quinqueloculina seminulum* (Linné, 1758), *Hanzawaia* sp. cf. *H. tagaensis* (Asano), and *Criboelphidium* sp. All species are common to nearshore open marine, reefal, or lagoonal environments. Table E-6 documents the total population and the total and relative abundance by species.

Table E-7 summarizes the ecological parameters controlling the occurrence of the seven mollusc species documented in this study. Table E-8 presents the list of mollusc species identified in all three cores and their total population, as well as total and relative abundance by species. The only nonmarine species was *Physa virgata* (Gould, 1855); the marine species consisted of *Cerithidea* (*Cerithidopsis*) *pliculosa* (Menke, 1892), *Littoridina* (*Texadina*) *sphinctostoma* (Abbot and Ladd, 1951), *Polinices duplicatus* (Say, 1822), *Argopecten irradians* (Lamarck, 1819), *Chione latrilata* (Conrad, 1841), and *Callocardia* (*Agriopoma*) *texasiana* (Dall, 1892).

The Ostracode Record

The ostracode population recorded at Cayo del Oso Bay indicates the basin received both seawater and freshwater overtime. Close to site 41NU2, core 2 contained no marine ostracodes. However, three specimens of the foraminifer *A. beccarii* intruded into the area suggesting eolian transport. *C. glaucus*, *C. rawsoni*, and *Cypridopsis* sp. rarely entered the basin. The taphonomic parameters indicate the specimens were introduced by streamflow and were unable to settle and develop a biocenosis (Figure E-5). Fragmentation and abrasion were moderately high (15-30%). The poor record is also an indication of a reworked fauna (Table E-2).

Core 4 contained a diverse assemblage mostly of marine species. A brackish-water ostracode alternated with the marine forms. *C. salebrosa*, a limnic to upper bay species, occurred throughout the core ranging from extremely rare (1) to common (22). Occasionally, it was the only species present in the sediments. The upper- and mid-bay ostracodes present at core 4 were *P. rugata*, *R. wantlandi*, *A. laevicula*, *C. sandbergi*, *L. avellana*, and more rarely *X. banda*, and *H. probosciduala*. In general, the ostracode population at core 4 represented a biocenosis (Figure E-6). Fragmentation and abrasion ranged from low to moderately low (5-15%), occasionally moderately high (30%). Some shells showed evidence of coating (5%) indicating shallow water and wave action (Table E-2). Both the marine and limnic populations established communities as indicated by the disarticulation and adulthood ratios (Table E-4).

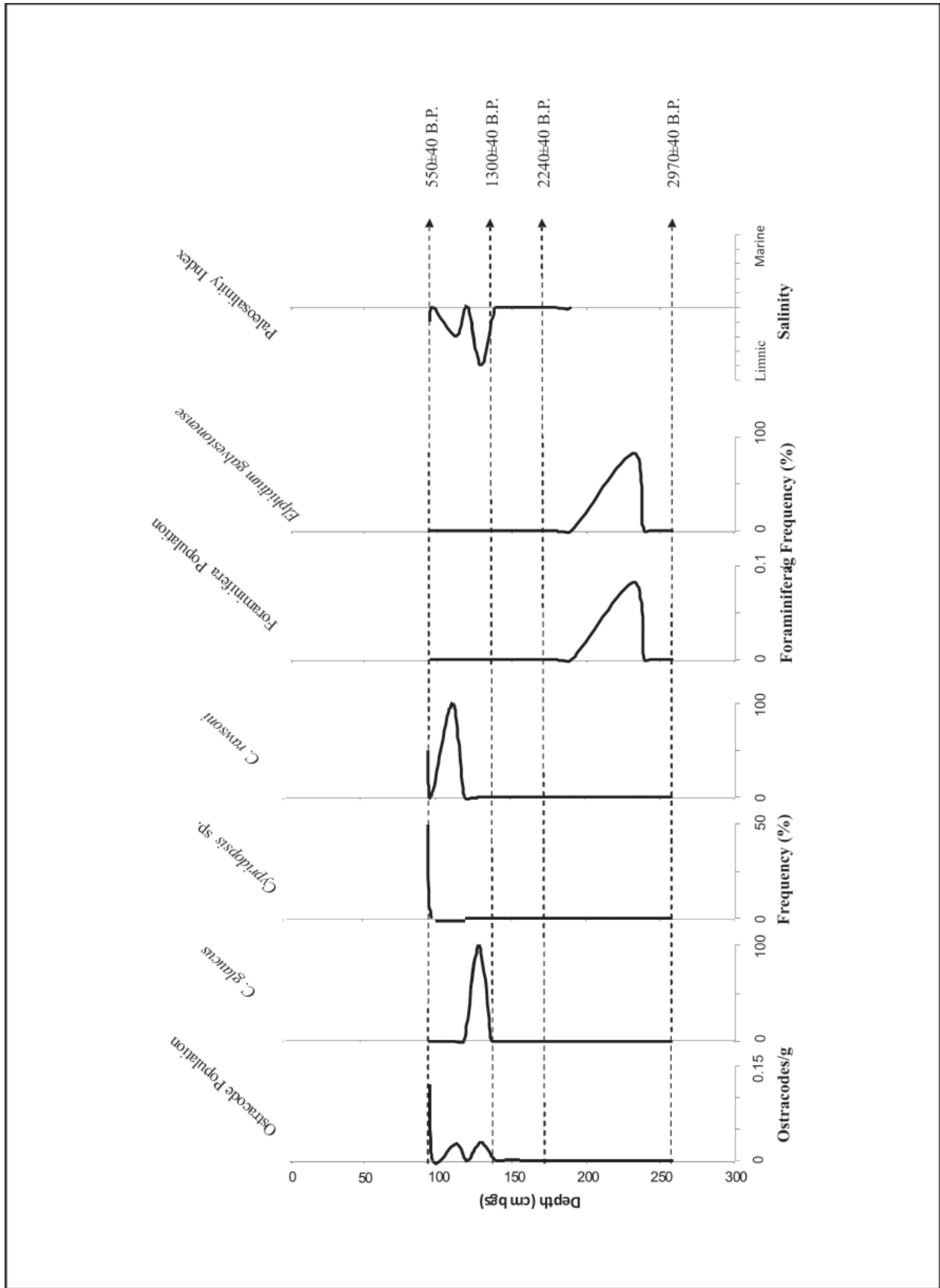


Figure E-5. Core 2 ostracode and foraminifera population total and relative abundance with core paleosalinity index diagram (extreme right).

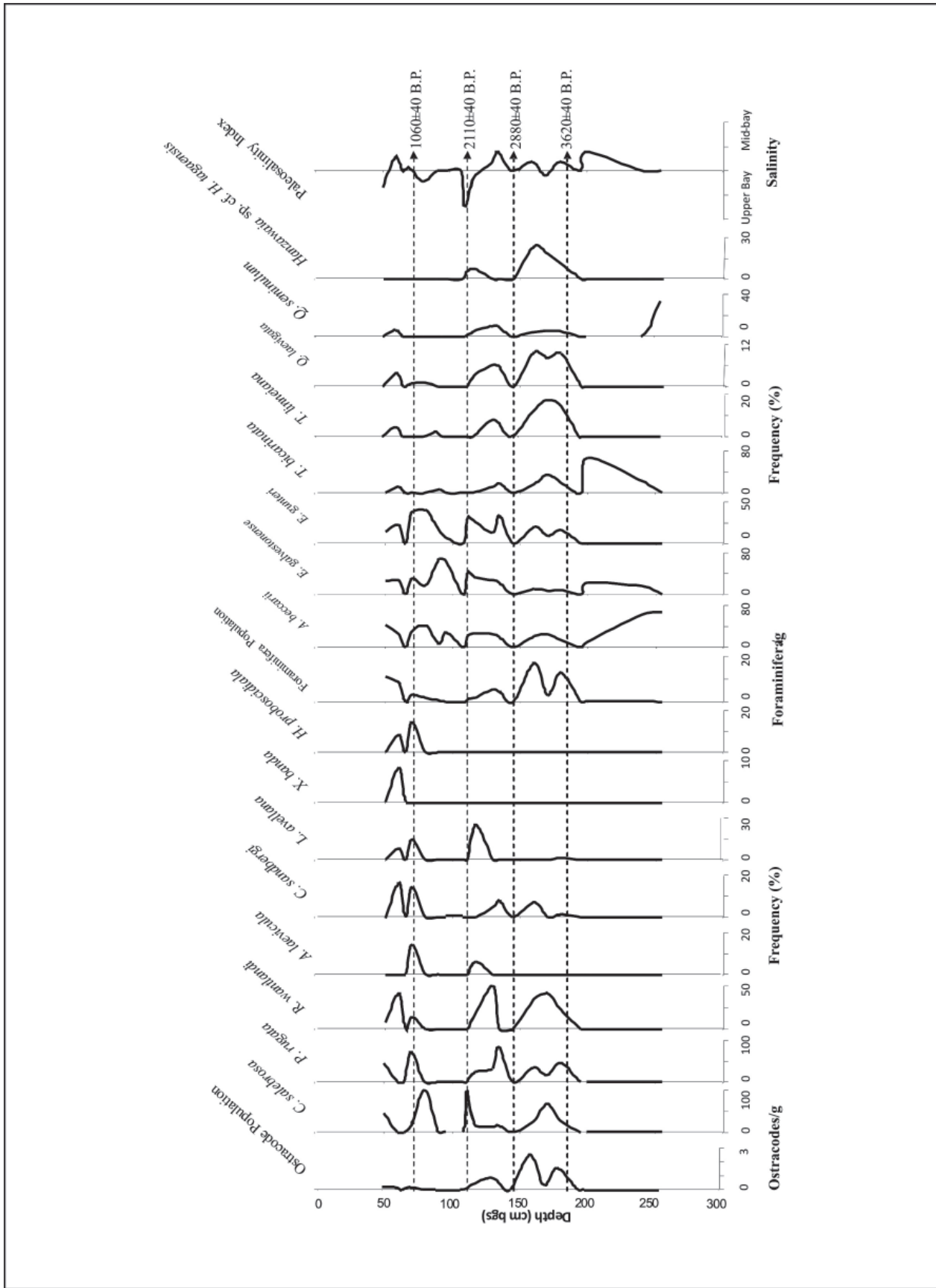


Figure E-6. Core 4 ostracode and foraminifera population total and relative abundance with core paleosalinity index diagram (extreme right).

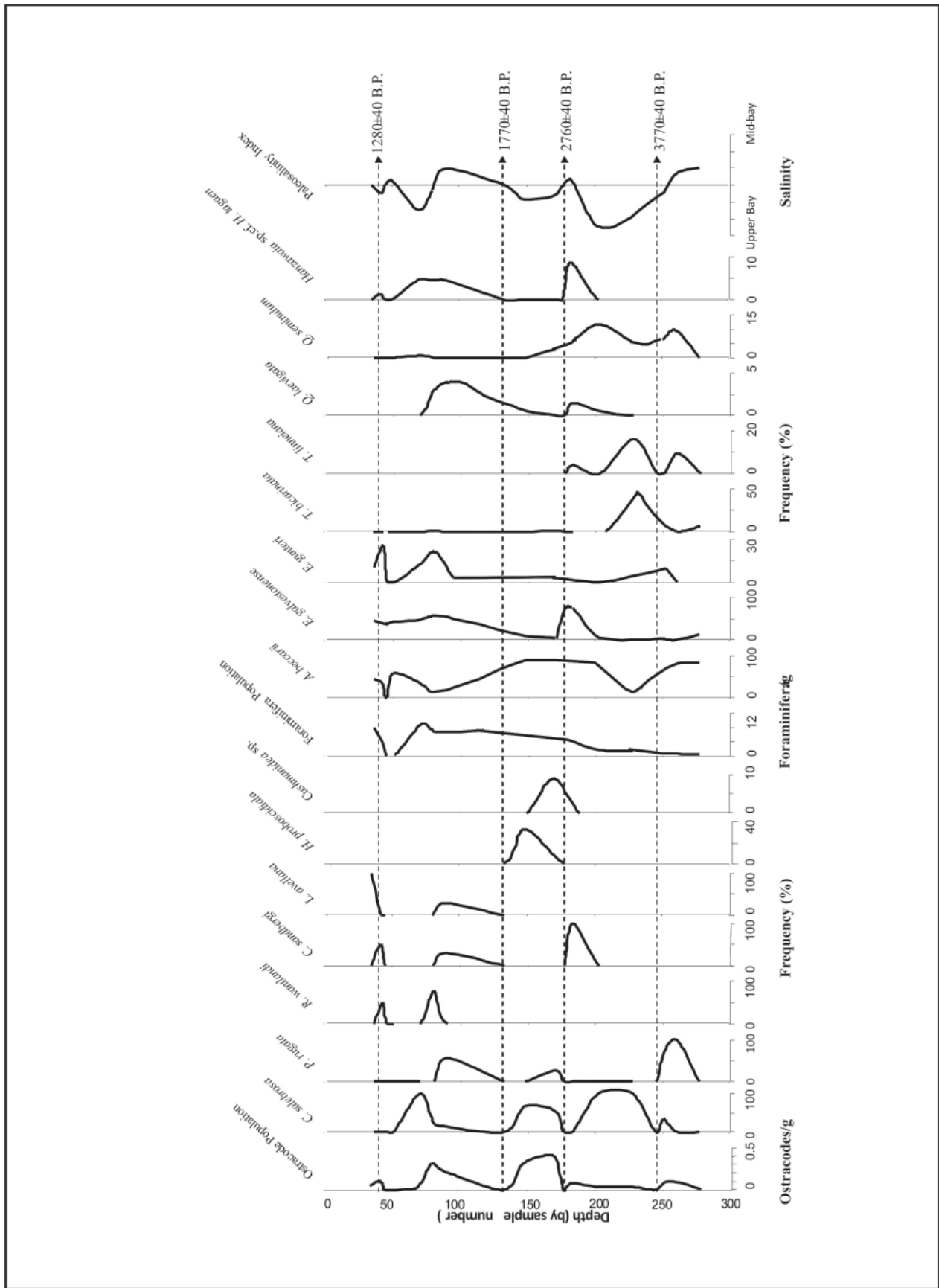


Figure E-7. Core 6 ostracode and foraminifera population total and relative abundance with core paleosalinity index diagram (extreme right).

Table E-5. Ecological Requirements of Foraminifera Species Recovered from Site 41NU2, Oso Bay, Nueces County, Texas

Species	Habitat
<i>Ammonia beccarii</i> (Linné)	Nearshore or lagoonal
<i>Elphidium galvestonense</i> Kornfeld	Nearshore or lagoonal
<i>Elphidium gunteri</i> Cole, 1931	Nearshore or lagoonal
<i>Hanzawaia</i> sp. cf. <i>H. tagaensis</i> (Asano)	Nearshore or lagoonal
<i>Triloculina bicarinata</i> (d'Orbigny)	Reef or lagoonal
<i>Triloculina linneiana</i> (Brady, 1884)	Reef or lagoonal
<i>Quinqueloculina laevigata</i> (d'Orbigny)	Reef or lagoonal
<i>Quinqueloculina seminulum</i> (Linné 1758)	Reef or lagoonal

Core 6 contained a diverse fauna as well, also consisting mainly of marine species. At this location, *C. salebrosa* was less abundant (1-6 specimens) suggesting higher seawater influence. However, its occurrence alternated with the marine forms like at core 4. The alternating upper- and mid-bay species included *P. rugata*, *R. wantlandi*, *C. sandbergi*, *L. avellana*, *H. proboscoidiala*, and *Cushmanidea* sp. (Figure E-7). Low fragmentation and abrasion (<5%) characterized the specimens. No other taphonomic parameters were recorded. The ostracode population also established a biocenosis (see Table E-4), with the possible exception of *C. salebrosa* that might have been transported to the site from areas that received freshwater discharge. As a result, its occurrence at core 6 indicates pulses of freshwater input.

The Foraminifera Record

The foraminifera population recorded at Cayo del Oso Bay indicates the basin received a strong marine influence throughout the record. Restricted to marine environments, only a few species may tolerate brackish-water conditions with a salinity range greater than 10‰ (parts per thousand). As indicated earlier, core 2 contained no marine forms except for interval 129 where three specimens of *A. beccarii* were identified (Figure E-5). The poor foraminifera record in core 2 also indicates that these organisms were introduced during a storm. The taphonomic parameters show moderately high fragmentation and abrasion (15-30%) (Table E-2).

Core 4 contained a diverse and abundant population consisting of *A. beccarii*, *E. galvestonense*, *E. gunteri*, *T. bicarinata*, *T. linneiana*, *Q. laevigata*, *Q. seminulum*, and *Hanzawaia* sp. cf. *H. tagaensis* (Table E-5). The foraminiferal assemblage represented a biocenosis with a population size ranging from extremely rare (3 specimens) to very abundant (606 specimens) (Table E-2 and Figure E-6). Fragmentation and abrasion ranged from low to moderately low (5-15%),

occasionally moderately high (30%). Some tests showed evidence of coating (5%) consistent with shallow water and wave action. In general, *A. beccarii* (>25%), *E. galvestonense* (>15%), and *E. gunteri* (>10%) were the most common and abundant species, rarely overshadowed by *T. bicarinata*. Other species occurred randomly at lower concentrations (Table E-6).

Core 6 was equally diverse and abundant. The same eight species (Table E-5) occurred in the stratigraphic column. *E. galvestonense* was more abundant than *A. beccarii* in the upper part of the core (samples 295 to 300), but the latter was more abundant in the lower part of the core. *E. gunteri* was less abundant than in core 4 and the remaining species occurred in a patchy pattern (Table E-6). Low fragmentation and abrasion (<5%) characterized the specimens. No other taphonomic parameters were recorded. The foraminifera also established a biocenosis (Table E-4).

The Mollusc Record

By contrast with ostracodes and foraminifera, molluscs were less diverse and significantly rare. One freshwater gastropod and six marine gastropods and bivalves occurred in the Cayo del Oso cores. The freshwater snail *P. virgata* occurred once in core 2 (upper part). No other molluscs were recorded from core 2. The two specimens recovered showed evidence of fragmentation and abrasion suggesting they were introduced into the system.

The marine species that occurred in cores 4 and 6 only included: the gastropods *C. (C.) pliculosa*, *L. (T.) sphinctostoma*, *P. duplicatus*, and the bivalves *A. irradians*, *C. latirilata*, and *C. (A.) texasiana*. The mollusc population ranged from extremely rare (1 specimen) to rare (<17 specimens) (Table E-8). All six species were present in core 4 where *Cerithidea*

Table E-6. Foraminifera Composition Including Total and Relative Abundance of Species Present.

Sample ID	Depth (cm)	Total Depth (cm bgs)	Conventional ¹⁴ C Age* (Years B.P.)	Bulk Wt. (g)	Fraction Wt. (g)	Foraminifera	Foraminifera /g	Foraminifera	Δ beccarii	%	#	En. galvestonense	%	#	F. gunter	%	#	H. sp. cf. H. tagensis	%	#	T. bicarinata	%	#	T. linnei	%	#	Q. laevigata	%	#	Q. seminulum	%	#
118	93-96	95	550±40	32.9	5.1																											
67	96-98	97	550±40	36.7	4.5																											
119	4-8	112		28.8	2.0																											
120	14-17	120		33.8	1.2																											
121	23-26	129		36.6	2.5																											
122	32-35	138		37.8	3.1																											
69	36-38	140	1300±40	31.9	2.2																											
123	42-44	143		43.4	3.0																											
124	52-56	155	2240±40	47.3	2.8																											
125	62-64	163		54.7	4.9																											
126	70-73	170		30.4	4.3																											
127	80-82	180		36.8	5.1																											
128	90-93	190		51.5	7.2																											
129	24-26	232		606	17.3	606	11.33	258	42.6	25.7	156	42.6	14.9	90	14.9	2	0.3	21	6.8	1	2.3	4.4	18	3.8	29	6.1						
130	32-34	239		473	16.3	473	8.49	118	24.9	26.6	126	24.9	22.2	105	22.2	56	11.8	21	4.4	18	3.8	29	6.1									
131	40-42	244		33.1	9.2																											
132	49-51	252		47.1	16.7																											
133	56-58	259	2970±40	47.5	14.0																											
240	48-52	30		40.6	14.6																											
241	60-63	60		59.1	14.9																											
153	65-68	65	1060±40	38.8	6.2																											
243	80-83	80		25.5	4.0																											
244	88-90	90		18.5	3.1																											
245	95-97	95		34.4	7.4																											
156	13	108	2110±40	35.9	7.1																											
247	15-17	111		42.1	8.8																											
248	22-24	117		16.1	6.4																											
250	38-42	130		28.3	10.4																											
251	46-49	135		33.5	11.9																											
157	59-61	143	2880±40	19.2	7.3																											
252	75-78	160		34.4	7.4																											
253	85-88	170		42.1	8.8																											
254	98-100	180		16.1	6.4																											
158	14-16	195	3620±40	28.3	10.4																											
255	17-20	200		33.5	11.9																											
256	40-43	240		22.5	5.7																											
257	55-58	255		19.2	7.3																											
295	34-36	35		28.4	5.2																											
296	42-45	42	1280±40	34.7	5.5																											
299	44-45	45		39.0	4.0																											
297	50-53	50		32.5	5.1																											
298	67-70	70		18.9	5.1																											
299	77-80	80		27.2	7.9																											
300	87-90	90	1770±40	26.9	8.4																											
260	42-43	133		28.9	7.1																											
301	56-58	148		29.8	7.9																											
302	78-80	170		26.9	8.4																											
261	86-87	177	2760±40	28.9	7.1																											
303	91-93	183		29.8	7.9																											
304	9-13	203		35.8	8.1																											
305	22-25	228		32.9	8.4																											
262	40-42	245	3770±40	35.8	8.1																											
306	45-48	251		32.6	7.7																											
307	55-58	261		32.6	7.7																											
308	73-75	278		32.6	7.7																											

* Radiocarbon dates obtained from humate samples. For details see Tomka et al. (This volume)

(*C. pliculosa*) dominated throughout the stratigraphic column. *Callocardia (A.) texasiana* and *Chione latirilata* alternated with the former dominating the upper portion of the core, and the latter taking over the lower part of the core. Other species were extremely rare. Fragmentation and abrasion ranged from low to moderately low (5-15%), occasionally moderately high (30%). Some shells showed evidence of coating (5%) as for ostracodes and foraminifera, this pattern is consistent with shallow water and wave action. It is unclear if molluscs succeeded in establishing a community; it is possible, however, that at some intervals (e.g., 240, 247, 252, and 253) some species settled a biocenosis.

Core 6 contained a poorer mollusc record than core 4 (Table E-8). Three species randomly occurred at the site where the core was obtained: the gastropods *C. (C.) pliculosa* and *P. duplicatus*, and the bivalve *C. latirilata*. Single specimens of *C. latirilata* only occurred at the lower portion of the core (samples 302 to 308). Low fragmentation and abrasion (<5%) characterized the specimens. No other taphonomic parameters were recorded. Based upon the poor record it is unwarranted to infer if molluscs established a community.

Interpretation

Inferring environmental change in Cayo del Oso Bay (site 41NU2) required integrating the paleoecological signature of the micro-invertebrates found in the sediments. Ostracodes were the primary source for reconstructing the late Holocene history of the bay. The paleosalinity index generated from equation (1) was then compared with the foraminifera and mollusc signatures to establish the patterns of water salinization and dilution over time. The paleoenvironmental reconstruction of Cayo del Oso will be presented by core and summarized in the discussion.

Core 2. At the location of core 2 the particle-size analysis obtained from sample preparation shows that the area received a slow flow for the past 3000 years (Figure E-4). Sedimentation rates ranged from 0.02 to 0.05-cm yr⁻¹. Small pulses of coarser grain-size at the base (between 2970 and 2500 years B.P.) and top (between 700 years B.P. and the present) of the stratigraphic column, respectively, might be the result of freshwater discharge into the bay. Two episodes of slower discharge into the bay were recorded at about 2500 years B.P. and 1000 years B.P. when clay concentration exceeded 94%.

Biologically, core 2 was poor. All micro-invertebrates were extremely rare when present. Ostracodes, foraminifera, and molluscs indicate the fauna was transported into the site and

failed to settle a biocenosis. This interpretation is supported by the poor preservation of the shells and tests. A preliminary paleosalinity index (Figure E-5), however, indicates that prior to 2500 years B.P. a storm deposited extremely rare foraminifera (*E. galvestonense*). Later, over the past 1000 years, freshwater input into the basin introduced a few ostracodes from upstream. The three species present are euryhaline (wide-salinity tolerant) (Table E-3); thus capable of tolerating dilute brackish-water conditions. Contemporary to ostracodes, extremely rare gastropods (*P. virgata*) were introduced into the system as well. The paleoenvironmental inference of core 2 is limited and unwarranted.

Core 4. At the location of core 4 variable conditions were recorded by the particle-size analysis obtained from the sample preparation. Subject to sedimentation rates between 0.04 and 0.07-cm yr⁻¹, at least three pulses of coarser sediments characterized the stratigraphic column during the past 3620 years (Figure E-4). Around 3700 years B.P. and between 2110 and 1060 years B.P. the most conspicuous episodes of sediment inflow occurred at the location of core 4. The interval between approximately 3500 and 2110 years B.P. shows a period of slow flow accumulating clayey sediments (>74% clay). The rest of the core mostly consisted of sandy silt (Table E-1).

Biologically, core 4 contained a moderately diverse and rich faunal assemblage including foraminifera, ostracodes, and molluscs. Foraminifera were the most abundant microfossils in the stratigraphic column indicating the strong marine and marginal marine influence into Cayo del Oso Bay. Ostracodes and molluscs were not very abundant; diversity ranged from six to eight species most of them indicative of marine and marginal marine conditions. Amongst ostracodes, only *C. salebrosa* has limnic to upper bay affinities; hence it is interpreted as a brackish-water indicator associated with freshwater input. Its occurrence is consistent with the episodes of sediment coarsening approximately 3700 years B.P. and between 2110 and 1060 years B.P. During the interval of fine sedimentation upper bay to mid-bay ostracodes, foraminifera, and molluscs dominated the stratigraphic column (Tables E-4, E-6, and E-8).

Foraminifera were rare to very abundant. The estuarine species *A. beccarii* co-occurred with *E. galvestonense* and *E. gunteri* throughout most of the core. The porcelaneous forms of stronger marine affinities occurred in the lower part of the stratigraphic column (Figure E-6).

Ostracodes and foraminifera settled a biocenosis, as indicated by Tables E-4 and E-6. The taphonomic parameters show low to moderately high fragmentation and abrasion (5-30%); low

coating (5%) and rare encrustation indicating alternating conditions between fresh- and seawater input (Table E-2). The ostracodes adulthood and disarticulation ratios are consistent with the establishment of a community. In addition, foraminifers ranged significantly in size, suggesting that the different species successfully settled in the vicinity of core 4. The paleosalinity index derived from ostracodes shows three main events of water dilution consistent with the occurrence of *C. salebrosa* and increasing grain-size (Figure E-6). Foraminifera were in good agreement with the ostracode paleosalinity index.

Core 6. At the location of core 6, variable conditions prevailed in the site. Five well-defined coarse grain pulses were briefly interrupted by intervals of fine sediment accumulation with the exception of the interval between 1770 and 1280 years B.P. where clay-rich (>81%) sediments dominated the stratigraphic column. Four of the coarse-grain intervals occurred approximately between 4000 and 1770 years B.P. The last one was recorded more recently, after 1280 years B.P. (Table E-1 and Figure E-4).

Biologically, core 6 is similar to core 4. Ostracodes and foraminifers dominated; molluscs were rare to extremely rare. This association indicates a relatively stressful environment for ostracodes and molluscs. Foraminifera, however, thrived successfully in this environment. The ostracode *C. salebrosa*, was the only brackish-water species identified in the core. Upper bay ostracodes were not abundant either (extremely rare to rare) implying that salinity was greater than at core site 4. This interpretation is consistent with the abundance of foraminifers but not with the poor record of molluscs. By contrast with core 4, the taphonomic parameters indicate low fragmentation and abrasion (5%), the only features identified (Table E-2). In addition, the adulthood and disarticulation ratios of ostracode shells are consistent with a well-established biocenosis in the site (Table E-4).

A similar pattern is indicated by the abundant foraminifers that were composed of small and large specimens of the same species. Dominance of the estuarine species *A. beccarii*, associated with *E. galvestonense* and *E. gunteri* throughout most of the stratigraphic column indicates the constant input of freshwater into the system (Figure E-7). However, dominance of porcelaneous forms (e.g., *T. bicarinata*, *Q. laevigata*) in the lower part of the core suggests a stronger seawater influence prior to approximately 2000 years B.P. It is, however, at the interval between 3770 and 2760 years B.P. that the increasing population of *C. salebrosa* indicates the largest pulse of freshwater recorded at core 6. Three more freshwater pulses were identified later in the stratigraphic column.

Discussion

Understanding human occupation at site Cayo del Oso (41NU2) requires a thorough reconstruction of the lagoon's paleoenvironments. Based upon Cronin's (1986) semiquantitative classification of biofacies based on fossil assemblages found in six cores around Mustang Island and Garbett and Maddocks (1979) modern distribution data, it was possible to establish some patterns of environmental fluctuations over the past 4,000 years (Table E-9). This approach is particularly useful because previous records do not focus in this area but in a much broader region including most of the state (e.g., Bousman 1998; Caran 1998; Cooke 2005; Frederick 1998 Fredlund et al. 1998; Nordt et al. 1994, 2002). A few exceptions, however, include studies on sea level change during the Holocene with little detail to the Corpus Christi Bay area (e.g., Albert 2006; Blum et al. 2002; Ricklis 2004; Ricklis and Blum 1997; Ricklis and Cox 1998). Tomka et al. (2007) summarized the paleoclimate history of southern Texas based on soil organic carbonates, carbon isotopes, and arboreal and grass pollen obtained by several authors (Bousman 1998; Camper 1991; Cooke 2005; Nickels

Table E-9. Quaternary Ostracode Assemblages Used in this Study to Identify Environments of Deposition (after Cronin, 1986)

Depositional Environment	Salinity (ppt)	Ostracode Death Assemblage
Freshwater- fluvial	Limnic (0-0.5)	Marine and bay taxa absent; nonmarine taxa present include <i>Candona</i> , <i>Limnocythere</i> , <i>Physicypria</i> , <i>Ilyocypris</i> .
Upper bay	Oligohaline (>0.5-5)	Mixture of bay and freshwater taxa, including <i>Limnocythere</i> , <i>Candona</i> , <i>Perissocytheridea</i> , <i>Cyprideis salebrosa</i> , <i>Cytherura</i> .
Mid-bay	Mesohaline (>5-18)	Bay species only, including <i>Perissocytheridea</i> , <i>Cytherura</i> , <i>Megacythere repexa</i> , <i>Loxococoncha moralesi</i> .
Lower bay-inlet	Polyhaline (>18-30)	Mixture of bay and marine taxa, including <i>Actinocythereis subquadrata</i> , <i>Peratocytheridea bradyi</i> , <i>Proteoconcha</i> spp., <i>Cytherura</i> spp., <i>Hulingsina</i> spp., <i>Malzella littoralis</i> .
Sublittoral	Euhaline (>30-36)	Marine taxa only, including <i>Actinocythereis</i> , <i>Malzella</i> , <i>Pellucistoma</i> , <i>Puriana</i> , <i>Reticulocythereis</i> , <i>Hulingsina</i> , <i>Protocytheretta</i> .

and Mauldin 2001; Nordt et al. 2002). All of these data were collected around 200-300 km north from the gulf coast, hence providing a valuable but limited record.

In general, the region was subject to cooler, more mesic conditions during the Early Holocene (10,000-8000 B.P.) gradually drying towards the end of the 2000 year period. The Middle Holocene became drier and warmer (8000-4000 B.P.) briefly interrupted for relatively mesic conditions around 7000 to 6000 or 5000 years B.P., returning to xeric conditions between 5000 and 4000 years B.P. Paleoenvironmental proxy suggest that the Late Holocene (4000-0 B.P.) was cooler and/or wetter throughout most of the record, returning to more xeric conditions within the last few hundred years. Is this signature of the Late Holocene depicted by the ostracode and foraminifera record obtained from cores 2, 4, and 6 at Cayo del Oso Bay site (41NU2)?

The answer is not straightforward! However, the faunal trends show patterns that may be associated with cycles of fresh- and seawater input over time for the past 4000 years. Twelve radiocarbon dates, four from each core, provided the geochronological record of Cayo del Oso Bay. Core 2, the closest to site 41NU2, shows the youngest dates ranging from 2970±40 years B.P. (at 259 cm bgs) to 550±40 years B.P. (at 97-cm bgs). Two intermediate dates, 2240±40 years B.P. (at 160-cm bgs) and 1300±40 years B.P. (at 140-cm bgs), integrate the geochronometric record. Based on these dates, Core 2 indicates low to moderately low sedimentation rates ranging from 0.02 to 0.05-cm yr⁻¹ (Table E-1). The increasing sedimentation rate values corresponded to episodes of increasing grain-size that, in turn, corresponded with the introduction of nonmarine ostracodes during the latest episode (a limnic assemblage; Table E-9). No evidence of ostracodes was recorded before 1300 years ago suggesting the organisms did not settle or arrive to the site or did not preserve. The poor record and poor preservation of the few valves recovered implies the organisms were introduced into the pond but failed to settle a community. In addition, three tests of *A. beccarii* obtained towards the base of the core more likely resulted from eolian transport or overwash from the Gulf of Mexico during hurricanes. No further discussion of Core 2 is warranted, except that the occurrence of the continental water species *Cyprinotus glaucus*, *Candona rawsoni*, *Cypridopsis* sp., and *Limnocythere* sp. cf. *L. floridensis* indicates pulses of freshwater discharge from the small drainages to the west.

The paleoenvironmental history of Cores 4 and 6 differs from Core 2 but compare well with each other. Based on Table E-9 biofacies models two assemblages were recognized in Cores 4 and 6. Alternating occurrences of upper bay and mid-bay assemblages dominated the history of Cayo del Oso for

the past 4000 years. No indication of other assemblages was obtained from these cores. A comparison of the paleosalinity indices for each core generated for this study shows the possible alternating intervals of fresh- and seawater input into the area during the late Holocene (Figure E-8). Core 2 only tangentially correlates with the other two cores but it is incorporated in the reconstruction and discussion because of its geographic proximity to Cores 4 and 6 and its implications for human activity.

Tentatively, three paleoecological zones at Cayo del Oso were identified, each with minor fluctuations. From base to top, a mid-bay assemblage dominated the area before 3600 years B.P. Between 3600 years B.P. and 1300 years B.P. an interval of rapidly changing from mid-bay to upper bay assemblages characterized the site. The last 1300 years were characterized by an upper bay assemblage suggesting higher freshwater discharge into the basin.

The patterns shown by the paleosalinity indices of Cores 4 and 6 compare well with those obtained by Bousman (1998), Camper (1991), and Nickels and Mauldin (2001) using grass pollen from Patschke, Boriak, and Weakly Bogs (300 km north of 41NU2). Prior to 3600 years B.P. xeric conditions prevailed in the area, less precipitation and runoff occurred in Cayo del Oso Bay. A mid-bay ostracode assemblage dominated and porcelaneous foraminifera were common. Between 3600 and 2800 years B.P. more mesic conditions affected the area as shown by the occurrence of upper bay ostracodes and the dominance of the foraminifer *A. beccarii*.

From 2800 to 2100 years B.P. alternating upper bay and mid-bay assemblages and foraminifera indicate rapidly fluctuating, from mesic to xeric to mesic, conditions. This pattern is partially in good agreement with Robinson's (1982:598) reconstruction of Choke Canyon using phytoliths. However, Robinson's poor dating and underreported data make difficult to compare results.

Between 2100 and 1700 years B.P., increasing mesic conditions are indicated by the dominance of an upper bay assemblage and the sharp decline in foraminifera that slowly recovered during this period. The interval from 1700 to 1000 years B.P. experienced a sharp decline in ostracodes. Only *C. salebrosa* was recorded in this interval in Core 4; however, at Core 6 *P. rugata*, *R. wantlandi*, *C. sandbergi* and *L. avellana* accompanied *C. salebrosa* suggesting an increasing mid-bay influence. Foraminifera continued its slow recovery. In other words, the time interval of 1700 to 1000 years B.P. gradually turned more xeric, but variable as shown by the paleosalinity index of Core 6. The pattern shown by micro-invertebrates and foraminifera correlate well with the latest Holocene

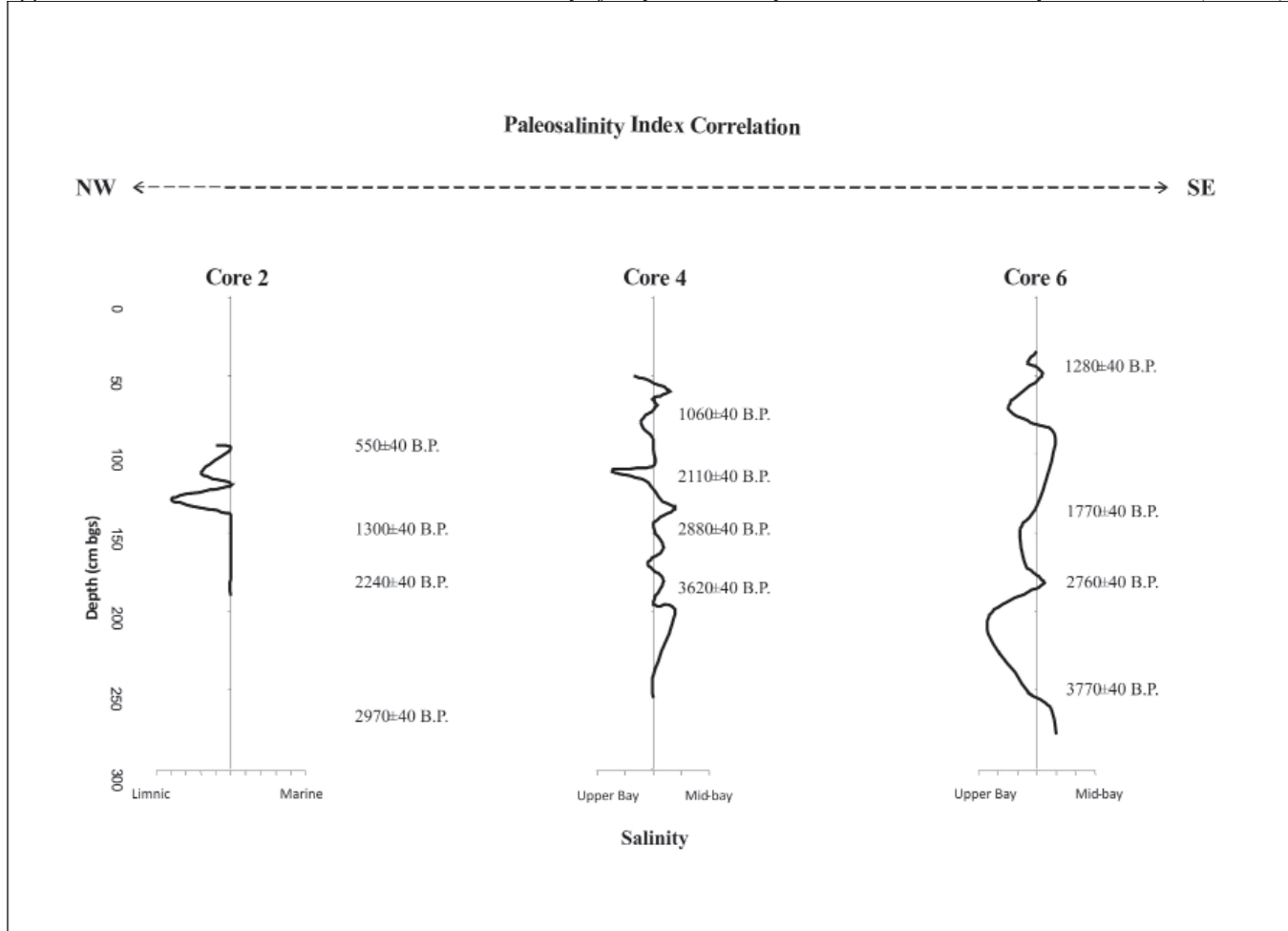


Figure E-8. Chronostratigraphic correlation of paleosalinity indices among the three cores from Cayo del Oso Bay, Nueces County, Texas.

grass pollen record at Patschke, Boriak, and Weakly bogs (Bousman 1998; Camper 1991; Nickels and Mauldin 2001).

1700-1200 years B.P.; <600 years B.P.) Ricklis and Blum (1997) found an increasing human population.

The last 1000 years consisted of an incomplete record both at Cores 4 and 6 that provided weak evidence of increasing moisture. Core 2 might reflect better the increasing precipitation introducing limnic species into the pond north of Cores 4 and 6. This pattern may correlate with the pollen records of Patschke, Boriak, and Weakly bogs (Bousman 1998; Camper 1991; Nickels and Mauldin 2001). However, ostracode abundance and diversity do not warrant the interpretation for the interval.

What is the impact of the environmental shift indicated by micro-invertebrates and foraminifera on human occupation? A comparison between the data generated for this study and that of Ricklis and Blum (1997; Table 1) suggests that humans encountered more favorable conditions during periods where the upper bay assemblage dominated in the environment. During xeric times (e.g., >3500 years B.P.; 2100-1700 years B.P.; around 1000 years B.P) human occupation was low; whereas during mesic intervals (e.g., 3500-2100 years B.P.;

Conclusions

In summary, micro-invertebrate and foraminifera analysis of the Cayo del Oso Bay site (41NU2) demonstrated the value of these microfossils as proxy tools for the reconstruction of marginal marine ecosystems over time. Based on ostracode diversity and abundance it was possible to recognize nonmarine from marine conditions during deposition of sediment in the region. Core 2 positively compared to cores 4 and 6, the combined signatures at the end of the record permitted a realistic, despite weak, reconstruction of the environment for the past 600 years. The close correlation between cores 4 and 6 confirmed field interpretations that both core sites would reflect similar conditions. However, minor variations between them were recognized that may be interpreted as micro-environmental variability.

The upper bay to mid-bay fluctuations recorded at Cayo del Oso Bay correlate well with pollen records from Central Texas

(Bousman 1998; Camper 1991; Nickels and Mauldin 2001) indicating alternating episodes of xeric to mesic conditions in the region over time. Micro-invertebrate and foraminifera data also agree with episodes of human occupation responding to periods of increasing precipitation and freshwater input into

the bay. Conversely, humans withdrew from the area during episodes of decreasing precipitation. The successful correlation between micro-invertebrates and the study of Ricklis and Blum (1997) on human population demonstrates the shifting environments prehistoric people faces in southern Texas.

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APPENDIX F:
Evaluating the Paleoclimatic Potential of Coastal Clay Dunes
at the Cayo del Oso Site

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Introduction

The Cayo del Oso site is situated at the confluence of Oso Bay and Corpus Christi Bay (see Chapter 2, this volume for geographic details). The geologic history of this region and geographic setting of the site are discussed in detail in Chapter 2 (this volume) and will not be repeated here except to define specific points relevant to the chapter. A small spit has formed across the mouth of Oso Bay by longshore transport of sediment within Corpus Christi during the late Holocene (Morton and Payne 1984; Withers and Chapman 1993). This spit is now detached from the mainland and known as Ward Island. Two inlets are present either side of Ward Island that permit water flow between Corpus Christi Bay and Oso Bay: Oso Inlet and Blind Oso Inlet (Williams and Kraus 1998). To the southeast of Ward Island is the main entrance to Oso Bay known as Oso Inlet. On the northwest side of Ward Island is a shallow, now largely dry mud flat that is locally referred to as the Blind Oso. The site (41NU2) is situated along the southeastern margin of the Beaumont Formation uplands adjacent to the Blind Oso. Along Corpus Christi Bay, both the Pleistocene upland and Ward Island are bordered by multiple linear offshore sand bars that Williams and Kraus (1998) determined have diminished in number over the last 60 years and slowly migrate offshore. Away from Corpus Christi Bay, the southeastern edge of the Pleistocene Beaumont uplands and Ward Island are bordered by an extensive clay wind-tidal mudflat that is now inundated by wind driven tides (Withers and Chapman 1993:383). Immediately west of Ward Island and the Blind Oso is a small drowned creek valley that is oriented roughly east-west. This unnamed drainage is relatively short and is incised into the Beaumont Formation (see USGS 1925).

Oso Creek has been a tributary of the Nueces River for at least the last glacial-interglacial cycle given that the lowest reach of the stream is incised at the contact between Encinal Peninsula (which is part of the Ingleside sand, a Pleistocene age barrier island associated with the Beaumont formation) and the Beaumont Formation (Shideler 1986:7). Ricklis (1993:8) notes that the Nueces River was incised as much as 60 m below modern sea level during the last glacial maximum, and it is likely that the Oso Creek valley was entrenched as well. Sheidler (1986) obtained a core from within Oso Bay, just off the southeast side of Ward Island (present site of Texas A & M University, Corpus Christi) and identified approximately 2-3 m of Holocene sediment resting on top of Pleistocene fluvial deltaic sediments (see in particular Sheidler 1986:29, Fig. 4). This core (core A) was peripheral to the center of the Oso Creek valley and probably does not accurately reflect the depth of incision, but implies that Oso Creek may not have been very deeply incised at the point it joins present day Corpus Christi Bay. Morton and Payne (1984) mapped the base of the Pleistocene Nueces incised valley and this map suggests that

the Oso Creek valley at the confluence with Corpus Christi bay was incised approximately 80 m below modern sea level. Rising sea level during the Holocene drowned the incised valleys of both streams before they could be completely filled with sediment during the Holocene.

Clay Dune Formation

Clay dunes, lunettes, or *lomas*, as they are locally known, are a form of source-border dune that form from the deflation of sand sized aggregates of mud from fringing mud flats. Were it not for the salt contained in the sediment, clay dunes would not form because it is the growth of salt crystals in the drying sediment that break up the mud cracks into the sand and silt sized aggregates that are capable of being entrained by wind. Clay dunes are frequently associated with playa lakes and pans where it is the presence of gypsum that facilitates clay pellet production and subsequent dune growth. But in coastal dunes it is the presence of halite in the salt water that promotes the formation of clay dunes. Coastal clay dunes typically form at the transition zone between wind-tidal flats and the uplands. As a form of eolian landform, clay dunes are rather anomalous because the clay aggregates, once moist, are sticky and adhere to one another, which results in the dunes being generally quite stable, or if they move they tend to migrate towards the sediment source rather than down wind, away from it.

The driving force in clay dune formation is clearly one of local hydrology (cf. Bowler 1973; 1983) but studies of some dunes (e.g. Holliday 1997:67; Frederick 1998) have concluded that because local hydrology responds to regional climate, the lithological record of such dunes may have regional paleoenvironmental significance. Indeed, in continental settings typical of pans and plays lakes, the growth of clay dunes is intimately tied to the hydrology of the pan, and the periods of dune growth are often linked to periods of arid climate. In the case of coastal clay dunes, the climatic linkage is considerably more complex as the sediment source areas are intermittently flooded and desiccated, mostly in response to wind and tidal influences.

Three specific factors affect the exposure and inundation of the Oso Bay mudflats: tides, long period tidal ranges, and wind. In the particular case of Oso Bay, the tidal range is very small, on the order of 20 cm (based upon the tidal range at the Naval Air Station-Corpus Christi, immediately southeast of Oso Bay; Williams and Kraus 1998:500) and Withers and Chapman (1993) note that tides alone have little influence on the exposure of the Blind Oso mudflats. In addition to the diurnal tides, Withers and Chapman (1993:394-385) note that there are long-period variations in the bay water level that

leads to generally high water in late May and late October, and periods of extremely low water levels in late February and late July. Given the micro-tidal nature of the bay, tides alone are not the major factor in immersion or emergence of the mud flat. Rather, wind driven tides appear to be most significant. Withers and Chapman (1993:383) note that in the May-November period, the prevailing winds are southerly or southeasterly and during this period water often covers the mudflats near the dunes and surrounding Ward Island. Conversely, the northerly winds associated with cold frontal passages in the December to February period push water away from the Blind Oso mud flats.

Clearly, the hydrology affecting coastal dunes is considerably more complex than that of continental pans, but process related observations on coastal clay dunes by W. Armstrong Price (1958; 1963) nevertheless led him to suggest that coastal clay dune deposition was climatically linked. In one paper, Huffman and Price (1949:120) note that the formation of clay dunes on the Texas coast is strongly related to climate:

“In south Texas, movement of the clay pellets begins in the month of March and continues until November when the growth of the dunes is arrested by winter precipitation and reduced evaporation. Their growth is most rapid during the drier summer months and during periods of prolonged drought.”

In another paper, Price and Kornicker (1961:249) noted that in the seven-year drought of the 1950s, some dunes accumulated a foot or more of fresh sediment. Price and Kornicker (1961:246) have also noted a climatic influence on the size of clay dunes with the dunes being highest near the Rio Grande Delta and lower toward the progressively more humid climates to the north and south.

Potential of Coastal Clay Dunes to Yield Palaeoenvironmental Records

Given these observations, it stands to reason that the periods of dune growth and development may provide a climatically sensitive record, whereby the dune grows during periods of relative aridity and become stabilized during more moist periods when soils form. In order to test the paleoclimatic utility of coastal clay dunes we performed a comparison of the paleoenvironmental records of two adjacent clay dunes at the Cayo del Oso site. The two dunes, the North Dune and the South Dune, are separated by a low relief ridge of the Pleistocene Beaumont Formation.

For this project we compared the stratigraphy of the two different dunes, and radiocarbon date the buried soil A

horizons. If climate is the driving force behind periods of dune aggradation, then the two dunes can be expected to yield comparable paleosol sequences (in terms of the number of buried soils and their radiocarbon age). Likewise, trends in dune texture and lithology (e.g. calcium carbonate content) are likely to be similar as well. If, however, there are significant differences in the soil stratigraphy between the dunes, then other, possibly local, factors other than climate may be critical in the sequence of growth and stability of these clay dunes. Furthermore, comparison of the dune lithological record with the pollen record from a pond that formed behind the North dune, may provide a useful test of the hypothesis as well.

Methods

Field Methods

An Eijkelkamp percussion drilling set was used to extract two adjacent cores from the north dune and the south dune adjacent to Ennis Joslin Road, and from a former pond that was situated at the southwest side of the intersection of Ennis Joslin Road and Ocean Drive. Cores were collected in 5 cm diameter PVC pipe that were labeled and capped following retrieval. At each locality, the two cores were collected approximately 1 m apart and coring was terminated upon reaching the Pleistocene Beaumont Formation. The north dune is represented by Core 4, which was collected from near the intersection of Ocean Drive and Ennis Joslin Road. The south dune, represented by Core 6, was cored along the south side of Ennis Joslin Road near its intersection with Del Oso Road.

Upon completion the cores were taken to the Center for Archaeological Research at the University of Texas at San Antonio where they were cut open, cleaned, described and sampled for various paleoenvironmental analyses (e.g. pollen, phytolith, diatom, ostracod, radiocarbon, bulk lithology and micromorphology; see Figures F-1 and F-2 for graphic depictions of the core stratigraphy and the location of samples collected). At each locality, the most complete core was chosen for detailed analysis and description. The lithology and color of the strata present in each core was recorded, but pedogenic attributes such as structure was not recorded as it was difficult to evaluate from core samples. Only rudimentary descriptions were made for deposits judged to be disturbed or fill and these deposits were not sampled.

It is important to consider that the cores were collected from the project right-of-way immediately adjacent to Ennis Joslin Road, and not from the summit of each dune. It is not known how much, if any, effect this placement may have on

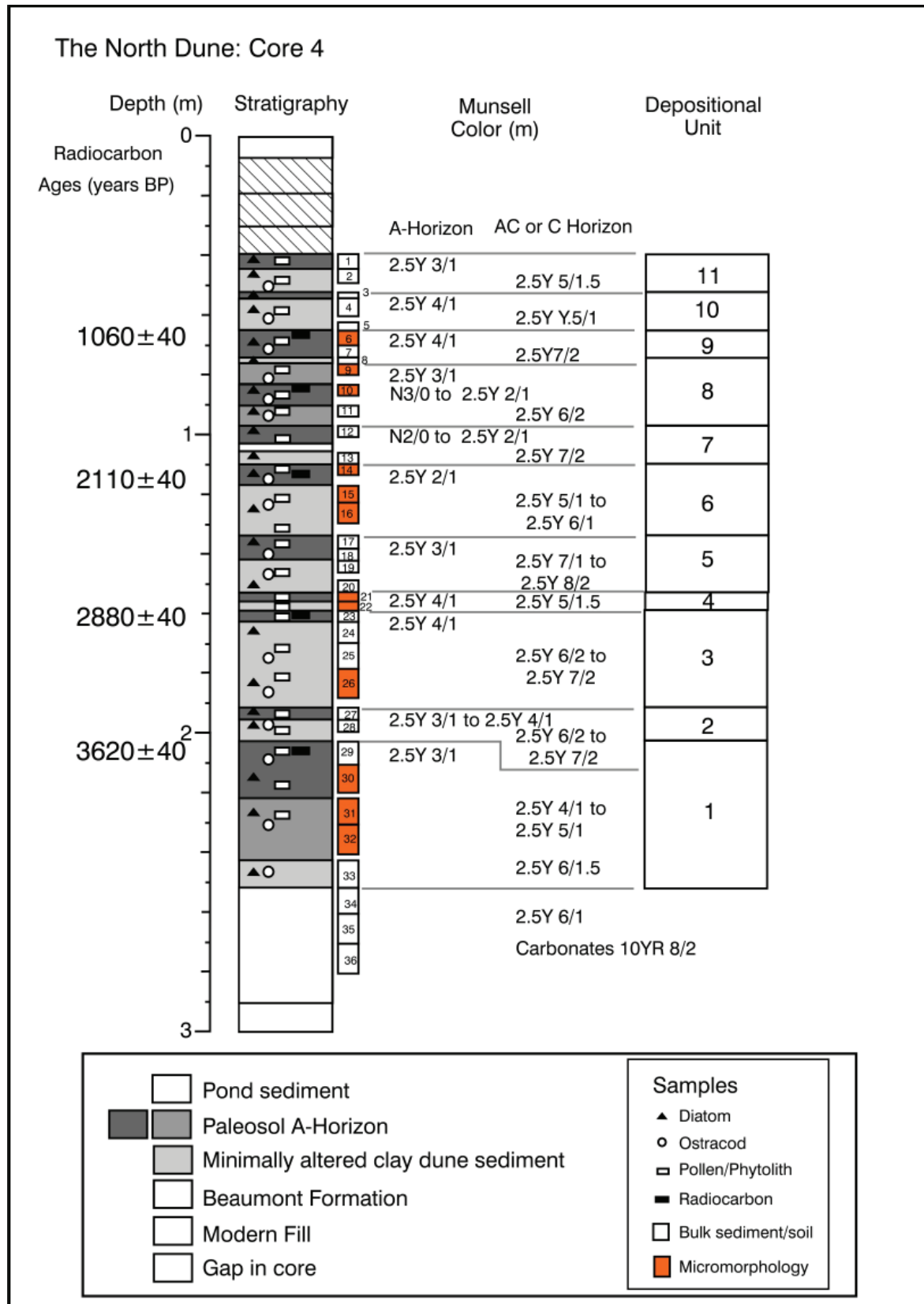


Figure F-1. An expanded drawing showing the core stratigraphy, Munsell color, and inferred depositional units in the North Dune core.

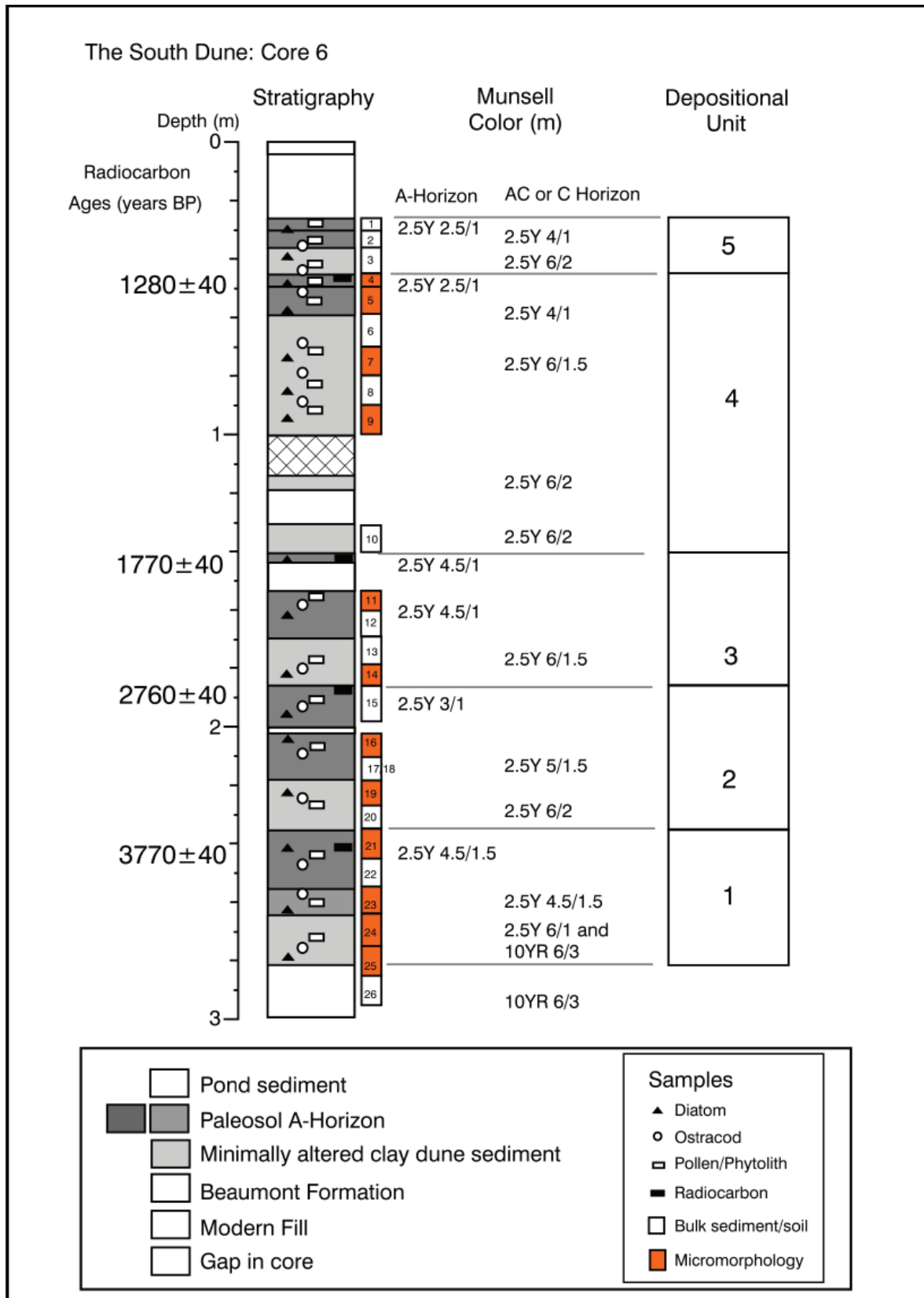


Figure F-2. An illustration of the core stratigraphy, Munsell color, and inferred depositional units in the South Dune core.

the stratigraphic results as there is little information in the literature on small scale variations in clay dune stratigraphy.

Lab Methods

A total of 82 bulk samples were analyzed from the three cores (36 from Core 4, 26 from Core 6, and 20 samples from pond deposits in core 2). For each sample the texture, organic carbon, calcium carbonate content, magnetic susceptibility, organic carbon content and stable carbon isotopic composition were determined. A select suite of samples was also collected for thin section preparation and soil micromorphological analysis. The details of the analytical methods employed are presented below, and the results of the lab work are presented in Tables F-1, F-2 and F-3.

Particle Size Analysis (Texture)

Texture analysis was performed using the hydrometer-sieve method (cf. ASTM 1985; Gee and Bauder 1986; Bouyoucos 1962). For this analysis the total sample was air dried and weighed, then crushed with a rubber pestle and mortar, and subsequently passed through a 2 mm sieve. Material caught on the 2 mm sieve, if detrital material, was then sieved at a 1 phi interval and the mass on each sieve recorded. A split of the <2mm size material (roughly 40 grams) was then soaked 50 ml of a 5% sodium hexametaphosphate solution overnight, and then mixed in a mechanical mixer for 5 minutes before being diluted to 1 liter with distilled water. This mixture was placed in a 1 liter settling jar, mechanically agitated for 1 minute, and then set on a table, after which point hydrometer readings were made a different time intervals (specifically 1,

Table F-1. Results of Lab Work Obtained for Samples Derived from Core 4

Sample No.	Rank	Depth	% $\geq 63\mu$	% $\geq 2\mu$	% gravel	% sand	% silt	% clay	Texture (SCS)	mean	sd	skew	kurt	χ	Organic C%	δ 13C	CCE %	Depositional Unit
205	1	42	33.35	57.77	0.00	33.35	24.42	42.23	clay	7.68	4.46	0.33	0.22	7.32	0.86	-15.67	6.1	11
206	2	47	31.26	55.60	0.00	31.26	24.34	44.40	clay	7.82	4.42	0.27	0.21	6.60	0.47	-15.03	10.3	11
207	3	53	31.89	60.22	0.00	31.89	28.33	39.78	clay	7.81	4.31	0.61	0.22	6.62	0.41	-15.09	4.6	10
208	4	57	29.01	53.97	0.00	29.01	24.96	46.03	clay	7.94	4.43	0.12	0.20	5.37	0.37	-15.05	13.7	10
209	5	64	28.83	61.56	0.00	28.83	32.73	38.44	clay	7.91	4.01	0.71	0.11	5.82	0.38	-15.43	12.8	10
210	6	67	25.54	61.35	0.00	25.54	35.81	38.65	clay	7.85	4.16	0.68	0.24	7.12	0.51	-15.49	4.0	9
211	7	72	29.45	58.87	0.00	29.45	29.42	41.13	clay	7.84	4.29	0.62	0.22	6.16	0.50	-16.15	6.7	9
212	8	75	33.35	81.95	0.00	33.35	48.60	18.05	loam	6.34	2.90	0.77	0.68	5.88	0.38	-15.61	7.6	9
213	9	78	31.3	66.10	0.00	31.30	34.80	33.90	clay	7.67	4.14	0.68	0.25	7.26	0.50	-16.56	2.4	8
214	10	85	34.4	62.17	0.00	34.40	27.77	37.83	clay	7.66	4.35	0.62	0.23	8.47	0.77	-15.21	2.1	8
215	11	92	37.02	61.27	0.00	37.02	24.25	38.73	clay	7.60	4.36	0.63	0.23	6.44	0.45	-15.03	7.0	8
216	12	98	32.75	58.37	0.00	32.75	25.62	41.63	clay	7.83	4.40	0.60	0.22	8.93	1.02	-16.67	2.1	7
217	13	108	28.28	57.93	0.00	28.28	29.65	42.07	clay	7.90	4.30	0.47	0.22	6.54	0.41	-16.75	12.5	7
218	14	112	29.07	59.35	0.00	29.07	30.28	40.65	clay	7.85	4.31	0.65	0.23	9.21	0.63	-15.64	4.0	6
219	15	120	22.36	49.06	0.00	22.36	26.70	50.94	clay	8.09	4.40	-0.23	0.03	7.40	0.57	-15.36	16.1	6
220	16	126	17.79	44.18	0.00	17.79	26.39	55.82	clay	8.27	4.39	-0.35	0.20	5.88	0.53	-15.87	28.0	6
221	17	136	16.12	43.51	0.00	16.12	27.39	56.49	clay	8.33	4.33	-0.34	0.21	5.70	0.50	-15.86	23.4	5
222	18	140	15.34	42.14	0.00	15.34	26.80	57.86	clay	8.35	4.32	-0.35	0.21	5.98	0.52	-16.56	19.8	5
223	19	144	14.44	39.63	0.00	14.44	25.19	60.37	clay	8.45	4.33	-0.40	0.21	5.98	0.50	-15.78	20.4	5
224	20	150	14.06	44.88	0.00	14.06	30.82	55.12	clay	8.30	4.24	-0.26	0.23	5.77	0.46	-15.95	23.1	5
225	21	154.5	21.69	48.94	0.00	21.69	27.25	51.06	clay	8.05	4.41	-0.11	0.22	9.44	0.61	-15.76	13.1	4
226	22	157.5	22.05	53.79	0.00	22.05	31.74	46.21	clay	7.95	4.29	0.25	0.23	5.09	0.46	-16.74	23.1	4
227	23	162	22	49.68	0.00	22.00	27.68	50.32	clay	8.33	4.17	-0.17	-0.02	6.25	0.48	-16.30	7.9	3
228	24	166	16.97	41.01	0.00	16.97	24.04	58.99	clay	8.37	4.45	-0.44	0.20	5.43	0.40	-16.01	17.9	3
229	25	174	18.29	43.28	0.00	18.29	24.99	56.72	clay	8.28	4.45	-0.39	0.20	5.89	0.38	-15.68	13.7	3
230	26	184	18.59	44.47	0.00	18.59	25.88	55.53	clay	8.25	4.44	-0.37	0.21	5.64	0.35	-15.49	13.7	3
231	27	194	18.18	48.81	0.00	18.18	30.63	51.19	clay	8.18	4.30	-0.22	0.21	7.32	0.58	-17.32	2.7	2
232	28	198	17.11	43.45	0.00	17.11	26.34	56.55	clay	8.28	4.37	-0.34	0.22	5.91	0.45	-15.23	7.9	2
233	29	216	36.51	70.06	0.00	36.51	33.55	29.94	clay loam	7.30	3.95	0.76	0.31	6.47	0.29	-17.30	0.6	1
234	30	224	35.18	65.53	0.00	35.18	30.35	34.47	clay loam	7.60	4.17	0.64	0.25	7.39	0.24	-17.66	0.0	1

Table F-1. Continued...

Sample No.	Rank	Depth	% ≥63μ	% ≥2μ	% gravel	% sand	% silt	% clay	Texture (SCS)	mean	sd	skew	kurt	χ	Organic C%	δ 13C	CCE %	Depositional Unit
235	31	236	37.52	66.08	0.00	37.52	28.56	33.92	clay loam	7.56	4.19	0.67	0.25	7.05	0.18	-20.01	0.3	1
236	32	245	38.25	67.96	0.00	38.25	29.71	32.04	clay loam	7.51	4.15	0.79	0.26	6.76	0.12	-18.28	0.0	1
237	33	256	44.41	69.95	0.00	44.41	25.54	30.05	clay loam	7.44	4.12	0.80	0.15	6.55	0.09	-18.21	0.6	1
238	34	266	51.30	76.39	0.00	51.30	25.09	23.61	sandy clay loam	7.03	3.80	0.80	0.35	5.24	0.04	-20.45	1.8	Beaumont
239	35	275	44.82	71.58	0.00	44.82	26.76	28.42	clay loam	7.01	3.80	0.76	0.36	5.72	0.07	-19.69	5.5	Beaumont
258	36	285	34.54	62.93	0.00	34.54	28.39	37.07	clay	7.71	4.32	0.62	0.23	6.09	0.12	-18.50	9.1	Beaumont

Table F-2. Results of Lab Work Obtained for Samples Derived from Core 6

Sample No.	Rank	Depth	% ≥63μ	% ≥2μ	% gravel	% sand	% silt	% clay	Texture (SCS)	mean	sd	skew	kurt	χ	Organic C%	δ 13C	CCE %	Depositional Unit
325	1	28	38.97	67.11	0	38.97	28.14	32.89	clay loam	7.50	4.23	0.66	0.26	7.20	0.60	-15.34	0.91	5
326	2	33	33.62	59.95	0	33.62	26.33	40.05	clay	7.79	4.38	0.63	0.22	6.46	0.42	-14.88	3.64	5
327	3	40	39.49	65.31	0	39.49	25.82	34.69	clay loam	7.62	4.23	0.79	0.24	6.16	0.35	-14.98	5.47	5
328	4	48	23.93	63.87	0	23.93	39.94	36.13	clay loam	7.79	4.05	0.69	0.25	8.33	0.64	-15.08	1.52	4
329	5	54	19.77	58.91	0	19.77	39.14	41.09	clay	8.20	3.96	0.58	0.03	6.56	0.48	-17.39	9.42	4
330	6	64	19.15	50.93	0	19.15	31.78	49.07	clay	8.14	4.31	-0.15	0.22	6.21	0.34	-17.12	11.84	4
331	7	75	15.69	48.79	0	15.69	33.10	51.21	clay	8.27	4.25	-0.22	0.24	6.24	0.30	-16.16	13.67	4
332	8	85	16.85	48.46	0	16.85	31.61	51.54	clay	8.23	4.29	-0.24	0.23	6.41	0.30	-15.55	10.93	4
333	9	95	17.61	50.33	0	17.61	32.72	49.67	clay	8.13	4.23	-0.19	0.23	6.33	0.28	-15.29	9.42	4
334	10	135	19.55	60.99	0	19.55	41.44	39.01	clay	7.90	4.04	0.13	0.25	6.48	0.30	-15.97	5.16	4
335	11	156	34.03	69.04	0	34.03	35.01	30.96	clay loam	7.27	3.99	0.53	0.31	7.20	0.38	-16.33	1.82	3
336	12	164	36.11	68.82	0	36.11	32.71	31.18	clay loam	7.30	4.11	0.56	0.29	7.24	0.36	-16.75	7.29	3
337	13	174	33.94	64.36	0	33.94	30.42	35.64	clay	7.46	4.30	0.38	0.25	6.93	0.24	-15.01	9.11	3
338	14	182	34.01	66.34	0	34.01	32.33	33.66	clay loam	7.37	4.21	0.38	0.29	6.43	0.22	-15.66	9.11	3
339	15	188	31.84	69.23	0	31.84	37.39	30.77	clay loam	7.41	3.97	0.58	0.17	10.22	0.37	-17.31	4.56	2
340	16	194	32.33	69.03	0	32.33	36.70	30.97	clay loam	7.44	3.96	0.59	0.31	9.58	0.33	-18.49	6.68	2
341	17	206	27.32	63.47	0	27.32	36.15	36.53	clay loam	7.66	4.09	0.30	0.27	7.82	0.33	-18.20	14.58	2
342	18	214	28.11	65.29	0	28.11	37.18	34.71	clay loam	7.61	4.05	0.33	0.27	7.64	0.28	-17.55	12.76	2
343	19	222	28.48	65.90	0	28.48	37.42	34.10	clay loam	7.57	4.02	0.27	0.28	6.76	0.27	-17.16	14.88	2
344	20	230	31.77	69.96	0	31.77	38.19	30.04	clay loam	7.40	3.93	0.58	0.32	6.91	0.22	-17.58	13.36	2
345	21	240	28.23	69.90	0	28.23	41.67	30.10	clay loam	7.44	3.86	0.53	0.32	9.29	0.28	-16.87	6.68	1
346	22	250	27.51	72.85	0	27.51	45.34	27.15	clay loam	7.29	3.71	0.59	0.36	9.95	0.34	-17.13	4.25	1
347	23	260	26.62	71.87	0	26.62	45.25	28.13	clay loam	7.54	3.60	0.68	0.23	10.16	0.26	-18.96	5.77	1
348	24	270	27.14	66.60	0	27.14	39.46	33.40	clay loam	7.63	4.04	0.61	0.28	9.34	0.15	-18.68	9.72	1
349	25	280	22.64	60.34	0	22.64	37.70	39.66	clay	7.85	4.17	0.17	0.25	8.47	0.21	-24.95	16.10	1
350	26	290	23.56	58.33	0	23.56	34.77	41.67	clay	7.89	4.23	0.10	0.23	8.08	0.10	-18.36	18.53	Beaumont

Table F-3. Results of Lab Work Obtained for Samples Derived from Core 2

Sample No.	Rank	Depth	% ≥63μ	% ≥2μ	% gravel	% sand	% silt	% clay	Texture (SCS)	mean	sd	skew	kurt	χ	Organic C%	δ 13C	CCE %	Depositional Unit
132	1	97	25.27	55.38	0	25.27	30.11	44.62	clay	7.87	4.48	0.04	0.24	5.89	na	na	2.13	Pond
133	2	108	12.45	45.95	0	12.45	33.50	54.05	clay	8.40	4.24	-0.32	0.21	6.55	2.03	-15.53	1.21	Pond
134	3	118	7.63	38.79	0	7.63	31.16	61.21	clay	8.77	4.05	-0.38	0.22	7.00	1.73	-15.59	1.21	Pond
135	4	128	6.41	37.63	0	6.41	31.22	62.37	clay	8.86	3.98	-0.38	0.23	7.17	1.58	-15.91	0.91	Pond
136	5	139	5.94	37.24	0	5.94	31.30	62.76	clay	9.22	3.62	-0.32	0.04	7.34	1.56	-15.61	0.91	Pond
137	6	146	6.53	34.93	0	6.53	28.40	65.07	clay	8.93	3.96	-0.40	0.24	7.31	1.42	-15.88	1.21	Pond

Table F-3. Continued...

Sample No.	Rank	Depth	% $\geq 63\mu$	% $\geq 2\mu$	% gravel	% sand	% silt	% clay	Texture (SCS)	mean	sd	skew	kurt	χ	Organic C%	$\delta^{13}\text{C}$	CCE %	Depositional Unit
138	7	156	7.36	36.60	0	7.36	29.24	63.40	clay	8.87	4.00	-0.40	0.23	7.44	1.42	-15.80	0.61	Pond
139	8	166	9.44	36.13	0	9.44	26.69	63.87	clay	8.83	4.04	-0.41	0.24	7.45	1.22	-15.77	1.21	Pond
140	9	175	9.08	34.79	0	9.08	25.71	65.21	clay	8.87	4.03	-0.42	0.24	7.55	1.18	-15.80	1.21	Pond
141	10	184	7.55	35.60	0	7.55	28.05	64.40	clay	8.89	3.97	-0.39	0.25	7.57	1.11	-15.59	1.21	Pond
142	11	195	9.01	39.76	0	9.01	30.75	60.24	clay	8.65	4.12	-0.36	0.21	7.45	0.98	-15.49	1.52	Pond
143	12	227	8.93	41.83	0	8.93	32.90	58.17	clay	8.57	4.18	-0.36	0.19	7.12	0.84	-15.44	1.52	Pond
144	13	235	14.74	48.40	0	14.74	33.66	51.60	clay	8.30	4.26	-0.26	0.18	6.94	0.56	-15.39	1.21	Pond
145	14	244	16.79	48.63	0	16.79	31.84	51.37	clay	8.27	4.31	-0.26	0.18	7.11	0.48	-15.70	1.52	Pond
146	15	253	16.38	42.87	0	16.38	26.49	57.13	clay	8.34	4.36	-0.37	0.05	7.45	0.38	-15.52	1.52	Pond
147	16	261	18.54	50.80	0	18.54	32.26	49.20	clay	8.14	4.28	-0.16	0.22	7.84	0.21	-16.78	1.52	Pond
148	17	269	31.64	59.48	0	31.64	27.84	40.52	clay	7.72	4.37	0.26	0.22	7.46	0.15	-18.82	8.50	Beaumont
149	18	277	38.1	67.57	0	38.1	29.47	32.43	clay loam	7.38	4.14	0.58	0.27	6.50	0.08	-21.73	16.10	Beaumont
150	19	286	39.35	70.96	0	39.35	31.61	29.04	clay loam	7.26	3.91	0.63	0.32	7.05	0.07	-21.54	17.01	Beaumont
151	20	295	39.21	71.61	0	39.21	32.40	28.39	clay loam	7.32	3.88	0.64	0.32	7.59	0.08	-24.60	13.36	Beaumont

3.5, 15, 45, 300, and 1440 minutes). A control hydrometer and temperature reading on an empty jar with nothing but distilled water and the sodium hexametaphosphate solution was made at intervals throughout the analysis to permit calibration of the hydrometer. A small split of the < 2 mm soil was also oven dried to determine the moisture content and correct the sample mass used in the hydrometer analysis (hygroscopic moisture correction). After 24 hours, the contents of the hydrometer jar were wet sieved through 37 micron sieve, and the sand retained on the sieve was transferred to a beaker and oven dried at 105°C. This sand was subsequently sieved at 0.5 phi intervals once dry and the mass retained on each sieve recorded. From these data the percentage of gravel, sand, silt and clay, as well as various descriptive statistics were calculated for the grain size distribution using a spreadsheet written by Paul Lehman.

Calcium Carbonate Content (C.C.E)

A 1.7 g split of the ground < 2 mm fraction of each soil sample was used to determine the calcium carbonate equivalent (C.C.E.) by means of a Chittick apparatus (Dreimanis 1962; Machette 1986). This sample was finely ground (to pass a 0.075 mm sieve), and then weighed, and placed into a small (250 ml) Erlenmeyer flask. Once attached to the Chittick apparatus, the liquid level in the measuring burette was set to -10 ml, then the stopcock was closed so no gas could leave the system, and the leveling bulb was dropped in order to establish a vacuum inside the sample chamber. At this point the barometric pressure and temperature in the room were recorded. Then 10 ml of 50% hydrochloric acid (ca. 6 N HCl) was delivered to the sample flask, which was agitated intermittently until the reaction had ceased

(usually 1-2 minutes). At this point, the leveling bulb was raised to the point that the liquid level inside of it was equal in elevation to the liquid in the burette, and the volume of gas evolved was then measured and the calcium carbonate equivalent calculated.

Organic Carbon

In order to determine the carbon content of each sample, splits of the < 2 mm size fraction were submitted to the Analytical Chemistry Lab at Institute of Ecology, University of Georgia. The carbon content was determined on a Micro-Dumas NA1500 Combustion Elemental (C/H/N) Analyzer (Carlo Erba Strumentazione, Milan). Details of the procedures used at in the procedure may be found on the institute's web page (<http://www.uga.edu/~sisbl/soilerb.html>) and general aspects of the method are discussed by Schulte and Hopkins (1996). Organic carbon content was determined from a 2 gram split of the < 2 mm size soil was treated with 6 N HCl to remove calcium carbonate, then rinsed repeatedly, subsequently dried at 105°C, and then finely ground once dry.

Stable Carbon Isotopic Composition

The stable carbon isotopic value of the organic carbon for each sample was determined at the Stable Isotope/Soil Biology Laboratory at Institute of Ecology, University of Georgia. These values were determined from the carbonate free < 2 mm soil used to determine the organic carbon content and were determined on a element ratio mass spectrometer by converting the organic carbon to a gas phase by extremely rapid and complete flash combustion of the sample material.

Magnetic Susceptibility

The reversible, low frequency mass magnetic susceptibility of the samples were determined by filling a 1 inch plastic cube with air dried <2mm size material from each sample. The samples were then weighed and then the low frequency magnetic susceptibility was measured on a Bartington MS2 meter and MS2B sensor. Each value was measured twice and the average values were used to calculate the reversible, low frequency mass susceptibility (χ_{lf}^c) which are reported in SI units ($10^{-8} \text{ m}^3 \text{ kg}^{-1}$). The precise methods and equations used may be found in Gale and Hoare (1991:222-226).

Petrographic Thin Sections

Small blocks of sediment cut from the bulk samples were submitted National Petrographic Inc., for thin section preparation, and were subsequently examined at a range of magnifications. Low magnification examination was performed with a Leica S8 APO binocular microscope fitted with transmitted light base and polarizing filters. Higher magnification examination employed a Leica DMEP polarizing light microscope.

Radiocarbon Dates

Bulk soil samples were collected from buried A horizons within each core and were submitted to Beta Analytic for radiocarbon assay. The organic carbon dated in these soils is largely pedogenic in nature and for this reason, the ages serve as minimum ages for the deposition of the unit they cap, and maximum ages for the next younger phase of eolian deposition.

Microfossils

In addition to the aforementioned basic physical analyses, samples from the two clay dunes were submitted to specialist for ostracode and diatom analysis. Winsborough (Appendix D; this volume) reports the results of the diatom analysis and, unfortunately, diatoms were rarely preserved in the clay dune deposits. Calcareous microfossils (ostracodes and foraminifera) were examined by Palacios-Fest (Appendix E; this volume). Although the preservation of these fossils was spotty as well, they were considerably more ubiquitous than the diatoms.

Results

General Observations

The dunes comprise approximately 2 to 3 m of alternating light and dark colored deposits which represent relatively

unaltered eolian sediment and soil A-horizons that formed during periods of relative stability and/or infrequent eolian sedimentation. The raw sediment (C horizons) ranges in color from light brownish gray (2.5Y6/2) to light gray (2.5Y 6/1 to 2.5Y 7/1), pale yellow (2.5Y 8/2) and gray (2.5Y 5/1.5). None of the deposits exhibited the pelletal structure they must have possessed at the time of deposition in either macroscopic or microscopic examination which suggests that these fabrics were relatively quickly changed by wetting and drying. The texture of the deposits ranged from clay to clay loam, with the deposits of the South Dune containing more silt and more often classified as clay loam rather than clay. The raw sediment appears to have contained between 0.3% and 0.4% organic carbon. The C horizons typically contained more calcareous microfossils than the pedogenically altered zones. This relationship is clear from Palacio-Fest's (Appendix E; this volume) counting and description of the foaminifera and ostracodes (see also Figure F-3 and Figure F-4) as well as the petrographic thin sections. Although the elevated calcium carbonate content within these zones is thought to be largely detrital, there was clear deposition of secondary calcium carbonate in the lowest soils in each core. The carbon isotopic composition of the pedogenically unaltered muds ranges from -18 to -24 in the earliest phase of deposition, and soon thereafter stabilizes between -15 to -17 per mil.

The thin A horizons (paleosols) ranged in color from black (N 2/0 and 2.5Y 2/1), to very dark gray (2.5Y 3/1) and gray (2.5Y 4/1), and generally were slightly coarser textured than the underlying C horizons. The coarser texture is expressed as a slightly greater sand content, and as a clear increase in the coarse: fine ratio observable in thin section. Nearly all of the A horizons exhibited significant calcium carbonate depletion. The leaching of calcium carbonate from these soils clearly degraded the carbonate microfossil content of these deposits, and is probably responsible for the slightly coarser texture of these zones. The carbon enrichment of the A horizons was quite variable, reaching a maximum value of around 1%, and a low of 0.3%, which is quite close to the detrital carbon content of the pedogenically unaltered sediment. Most of the soils formed in clay dune sediment exhibit slightly isotopically lighter values than the sediment within which they have formed, which may reflect the contribution of woody plants growing on the dunes during periods of relative stability. Although the isotopic composition of the soils are fairly similar, there is a subtle trend towards more C4 values throughout the sequence with early soils ranging around -17 per mil and the modern soils exhibiting values between -15 to -16 per mil (discussed in greater detail, later). Although the organic carbon content is one of the best indicators of pedogenic alteration of the deposits, the A horizons consistently exhibited very subtle

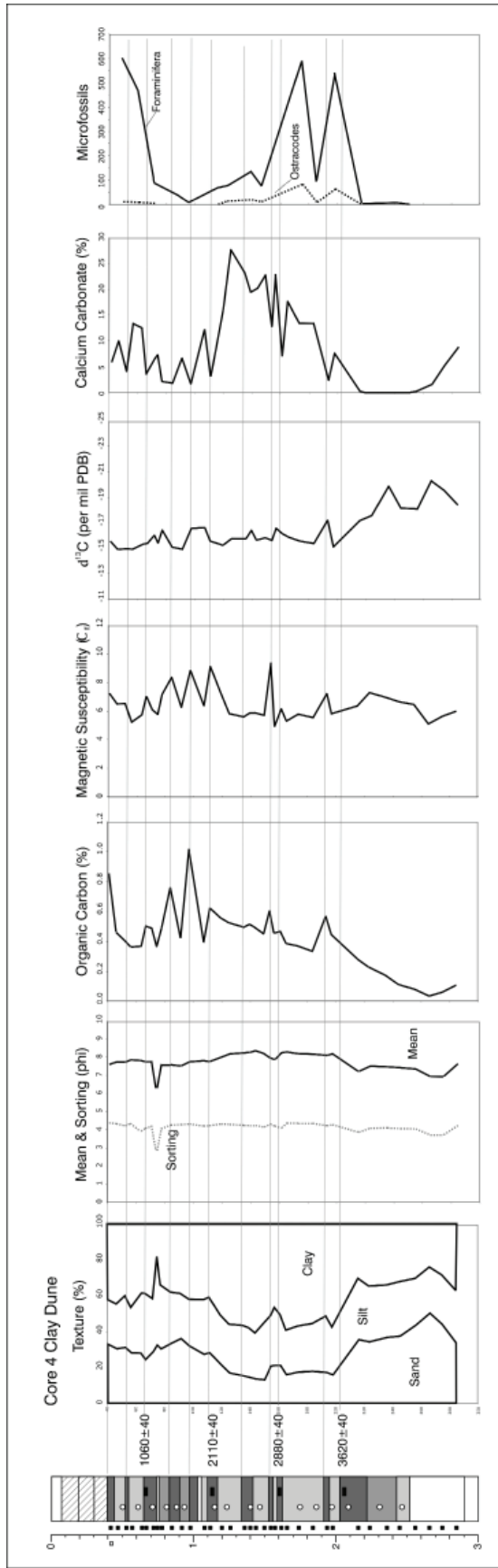


Figure F-3. Plot of the stratigraphy, radiocarbon dates, and physical properties of the North Dune core at the Cayo del Oso site. The light horizontal lines are registration lines for the top of buried soils and facilitate comparison of the different lithological attributes. Microfossil counts are from Palacios-Fest (this volume, Table 2).

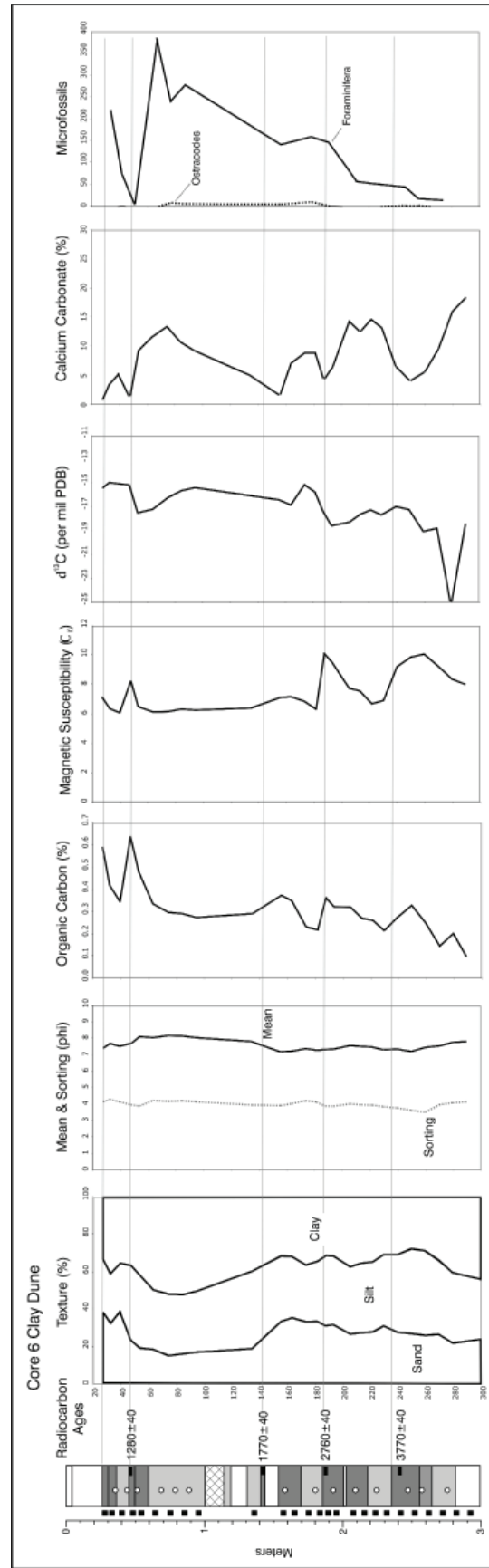


Figure F-4. Plot of the stratigraphy, radiocarbon dates, and physical properties of the South Dune core at the Cayo del Oso site. The light horizontal lines are registration lines for the top of buried soils and facilitate comparison of the different lithological attributes. Microfossil counts are from Palacios-Fest (this volume, Table 2).

but clear magnetic susceptibility enhancement. Most of the A horizons also exhibited local re-organization of the clay groundmass into cutans around framework grains and an optically active fabric, which was generally absent from the AC and C horizons.

A Comment on Chronology

As process observations by W. Armstrong Price (1958, 1963) and his colleagues (Price and Kornicker 1961, Huffman and Price 1949) demonstrate, the period of time represented by any one phase of clay dune aggradation may have been very short. The paleosols, therefore, undoubtedly represent most of the time in these stratigraphic records. Our chronology was obtained from radiocarbon ages obtained for bulk soil organic matter collected from the top of buried A horizons, and these dates appear to yield a solid chronology. But because of the nature of the depositional record, these dates do not reflect the time of deposition for the depositional unit from which they were collected. Rather, they most likely approximate the time of deposition for the next younger unit, and serve as a minimum age for the depositional unit from which they were collected

In hind sight, examination of the petrographic thin sections and the results of the granulometric analyses demonstrates that there is sufficient sand within the dunes to permit optically stimulated luminescence dating (OSL), and ideally, a paired sequence of OSL dates from the C horizons (dating the period of eolian sedimentation) and radiocarbon dates from the buried soils would yield the most useful geochronologic record. It is unclear at this time if the sand within the clay dune sediment was sufficiently bleached to accurately record the time of deposition, but this could, at least in part, be assessed if single grain OSL dates were employed.

The South Dune (Core 6)

The south dune is perched on the upland margin immediately north of and close to the mouth of an unnamed drainage. The core revealed a series of 5 discrete eolian depositional events that occurred in the last 4 thousand years (see Figure F-2) for a schematic illustration of the stratigraphy, samples collected, the deposit color and stratigraphic location of each depositional unit. In a broad sense, this sequence exhibits three significant textural cycles, with the first three and the fifth depositional events being similarly textured, and the fourth being considerably finer textured. The majority of the microfossils recovered by Palacios-Fest (Appendix E; this volume) were present within the fourth depositional unit. Figure F-4 charts the lithologic variation of the south dune deposits.

Although we do not have a date for the base of this dune, Ricklis (1997) obtained a radiocarbon age from an oyster

shell at the Beaumont-clay dune interface in this part of the site which yielded an age of 3,990±70 years B.P. (Beta-099304) which places the onset of dune sedimentation right around 4,000 radiocarbon years B.P. The first phase of eolian deposition was approximately 50 cm thick and a bulk soil radiocarbon date of 3,770±40 years B.P. (Beta-235810) from the top of the A horizon suggests that this phase was short lived, perhaps only 250 years or less. The stable carbon isotopes from the first depositional phase reflect a woodland environment at the start of the phase (with values of -24.95 per mil) which may be an artifact of deflation of the former Pleistocene surface soil. Carbon isotopic values gradually increase to -18.5 to -19 in the core of the deposit. The soil formed at the top of the first phase of deposition exhibited a $\delta^{13}\text{C}$ value of -16.87 per mil reflecting that relatively few C3 plants colonized the dune crest at this time. Few microfossils were found by Palacios-Fest (Appendix D, this volume) within the deposits of the first depositional phase, which may indicate that the lagoonal environment was not yet fully established when the dune began to form.

The second phase of deposition (DU2) was also around 50 cm thick and the soil at the top of this deposit yielded an age of 2,760±40 years B.P. (Beta-235809). It is probable that the deposition of this unit occurred closer to the 3770 years B.P. Sediments forming the core of this deposit exhibited $\delta^{13}\text{C}$ values between -17 to -18.5 per mil, and the soil at the top of the deposit yielded a value of -17.31 per mil.

A radiocarbon age of 1,770±40 years B.P. (Beta-235808) was obtained from soil at the top of the third depositional phase, which was around 40 cm thick. The fourth phase of dune aggradation was the thickest and finest textured in the entire sequence. A radiocarbon age obtained from the top of this deposit was 1280±40 years B.P. (Beta-235807) and indicates that more than a meter of sediment accumulated in less than 500 years. A sharp change in the $\delta^{13}\text{C}$ values in the soil at the top of this depositional phase suggests that the vegetation may have shifted towards C4 plants while the dune was stabilized as the top two samples differ by approximately 2.2 per mil.

The fifth and last phase of dune sedimentation is undated, and the soil formed in this deposit is presumed to be the modern surface soil. The deposit is only 25 cm thick and it is considerably coarser than the deposits of the fourth depositional unit.

The North Dune (Core 4)

The north dune overlooks the Blind Oso Inlet and core 4 penetrated almost 2.2 m of clay dune deposits that were found resting upon an eroded Beaumont Formation surface.

Although a similar lithologic record to the south dune was anticipated, the north dune revealed 11 depositional units (DU) and associated buried soils spanning approximately the same period of time as the south dune. Like the south dune, the north dune granulometry depicts three major cycles, starting with a relatively coarse phase, becoming finer textured between 3620 and 2110 years B.P., and becoming coarser again after 2110 years B.P. The following text describes trends plotted on Figure F-3.

The first phase of dune aggradation appears to be of similar age and magnitude as the first phase in the south dune. This deposit is approximately 50 cm thick and it is the thickest eolian deposit in the north dune record. A radiocarbon age from the top of this deposit yielded an age of 3620 ± 40 years B.P. (Beta-235806), and like depositional unit 1 in the south dune, Palacios-Fest found relatively few calcareous microfossils in this deposit. The $d^{13}C$ value of the buried A horizon found within this deposit was -17.30 per mil which is close to the value obtained for the soil in DU1 of the south dune.

Depositional Unit 2 was very thin (<20 cm) and consisted of an A-C soil. This deposit was not radiocarbon dated, and as such its age can be bracketed to the period between 3620 and 2880 years B.P. The subsequent phase of eolian sedimentation, DU3, was much thicker, and occurred sometime before 2880 years B.P. This deposit was approximately 30 cm thick and was capped by a very thin soil that was radiocarbon dated to 2880 ± 40 years B.P. (Beta-235805). The $d^{13}C$ values of the soil formed at the top of depositional unit 3 was -16.30 per mil. Samples from the C horizons of DU2 and DU3 contained numerous calcareous microfossils.

Three thin depositional units (4, 5 & 6) were deposited in the 770 year period between 2880 and 2110 years B.P. The calcium carbonate content of the raw sediment increases dramatically during this period, but somewhat surprisingly, the ubiquity of calcareous microfossils drops off. At the same time, the $d^{13}C$ values of the soils become progressively heavier, ending with a value of -15.64 per mil at the top of DU6 which was radiocarbon dated to 2110 ± 40 years B.P. (Beta-235804).

In the approximately 1000 year period between 2110 and 1060 years B.P. another three thin eolian deposits accumulated on the north dune, comprising a total of about 50 cm of new sediment. Judging by the color and organic carbon content of the sediment, these deposits were significantly altered by pedogenesis and as a result they yielded few calcareous microfossils. The first of these, DU7, was less than 20 cm thick and capped by a very black, organic rich soil which yielded a $d^{13}C$ value of -16.67, which is somewhat isotopically

lighter than the soils that immediately precede or follow it suggesting that the climate may have been somewhat more moist at this time. The following two units (DU8 and DU9) are fairly similar in character, and the top of DU9 was radiocarbon dated to 1060 ± 40 years B.P. (Beta-235803).

The last two depositional units in the core 4 record were deposited in the last millennium and it is during this period that stable carbon isotopic values reach their maximum (-15.49 per mil, in the soil at the top of DU10). Both of these depositional units comprise mostly unaltered sediment and very thin A horizons, and not surprisingly, the C horizons of these deposits yielded significant numbers of calcareous microfossils. The soil at the top of the sequence in DU11, the modern soil, yielded a $d^{13}C$ value of -15.67, which is slightly lighter than the soil in DU10, and this may reflect colonization of the dune by mesquite in the modern era.

Comparing the Two Records

General Observations

Even a cursory inspection of the two dune records (Figures F-3 and F-4) makes it clear that they are not reflecting the same sedimentation events and that some factor(s) other than climate must control dune sedimentation. None of the major lithological indices (texture, mean particle size, carbonate content) are consistent between the two dunes.

The textural composition of the dune deposits in both cores exhibit gradual, long term shifts which probably reflect changes in the nature of the mud flat adjacent to the dunes. Perversely, the lithology of the cores are out of synch with each other, with the core 4 record being clay dominated in the early period of dune formation, and becoming coarser textured in the later phases, whereas the core 6 record is just opposite of that, and is more coarse early on and becomes finer textured in the latter half of the record. Only the earliest and latest phases of deposition in each record appear to correlate, and one wonders if this would be the case if we had more accurate depositional ages for these sediments.

The microfossil work of Palacios-Fest (Appendix E; this volume) further supports this point. It was assumed that since these two dunes front the same mudflat (ostensibly the same sediment source), that the microfossil records contained within the dune sediment would be similar. But as Palacios-Fest has demonstrated, the two records are significantly different in both general trends (such as the paleosalinity index) and specific fauna present at any given time. It is clear that pedogenesis within the dunes is a major taphonomic factor controlling the preservation of ostracodes and foraminifera,

but even the most basic details such as the periods when the most microfossil rich sediment was deposited are out of sync between the two dunes.

The fact that the two records are not parallel demonstrates the significant role sediment source plays in the growth and development of these dunes, and how changes in dune texture are largely irrelevant in interpreting coastal clay dune history except as it pertains to changes in the sediment source immediately adjacent to the dune. It is clear that long-term changes in sediment size are not demonstrably climatic in nature. This finding is consistent with comments made by Ricklis (1997:17) that

“it is interesting and significant to note that a uniform stratigraphy across the area is not indicated. Rather, soil horizons tend to be horizontally localized, suggesting a complex history of clay dune formation, perhaps involving localized erosional events in which pre-existing sediments were partially removed prior to subsequent eolian deposition of new material.”

Hence, the hypothesis that periods of clay dune sedimentation and stability preserved within coastal clay dunes may be used as a climatic proxy must be rejected. As noted by Bowler (1973) the hydrology of the mud flat is the dominant control on dune sedimentation, and as this work has demonstrated, two dunes forming from the same mud flat may have dramatically different depositional records.

The dramatic stratigraphic variation exhibited by these two dunes (and the photos taken during the fieldwork as well) raise questions about the sedimentary and stratigraphic structure of coastal clay dunes. Most of the literature on these dunes document the physical processes of sedimentation, not the stratigraphy, and where we have stratigraphic records they are like the ones described here, made from vertical profiles. What is missing is a more comprehensive elucidation of stratigraphic variation of within a single dune, in three dimensions (e.g. stoss to lee, as well as longitudinally). Meissner (Figure 5-6; and personal communication 1997) observed a range of small-scale features within these clay dunes that reflect erosion and deposition of the dunes associated with runoff of water and subsequent gully formation, and these features are not presently represented in the coastal clay dune literature, and a more comprehensive look at a clay dune in the future may yield a more nuanced paleoenvironmental record.

Despite these problems, that does not mean that clay dunes do not hold paleoenvironmental potential.

Stable Carbon Isotopes

The stable carbon isotopic record from the dunes is a complex issue which needs careful consideration prior to interpretation. The raw records as plotted on Figures F-3 and F-4 contain two distinct carbon isotopic signals: 1) the composition of the detrital muds from the mud flat, and 2) the composition of the paleosol organic matter. Plotted together they show a complex back and forth pattern that appears to have little clear paleoenvironmental significance. But if the two signals are separated, a clear trend emerges. The upper panel of Figure F-5 plots the stable carbon isotopic ratio for soil organic matter in the top of the buried A horizons within the north and the south dunes. This plot uses the $\delta^{13}\text{C}$ values obtained by the University of Georgia lab, and not the values obtained by Beta Analytic for the radiocarbon ages. In the upper panel only soils for which radiocarbon ages are available have been used, with the exception of the uppermost soil which was assumed to be modern. It is clear that although the two records are relatively similar and depict similar trends, there are some differences in the soil carbon isotopic values at any one point in time. But the trend in both dunes is the same, namely a progressive shift towards a more arid adapted, C4 vegetation from 4000 to 1000 years B.P., and then a subtle shift back towards a slightly more C3 dominated mix in the recent period.

The lower panel of Figure F-5 plots show how the North dune record changes if all of the topsoil $\delta^{13}\text{C}$ values are plotted. This plot should be viewed with caution given that ages had to be assumed for soils for which radiocarbon ages are not available. The trend on this figure is essentially the same as the top panel with two notable exceptions. First, the anomalous value depositional unit 7, sometime after 2000 years B.P. The second is that the $\delta^{13}\text{C}$ maximum value is actually found in the soil of depositional unit 10, which dates sometime after 1060 years B.P.

From these plots it is apparent that the stable carbon isotopic record preserved by the buried soils within the clay dunes appears to preserve a record of gradually shifting vegetation conditions which are responding to climatic variation. The fortuitous preservation of numerous thin, short-term soils in the north dune appears to afford a reasonably well-resolved record of changing vegetation that, if were all radiocarbon dated may be useful in further understanding late Holocene climatic variation in this region.

Conclusions

The fundamental hypothesis tested by this study was that the lithological record of coastal clay dunes may be used as

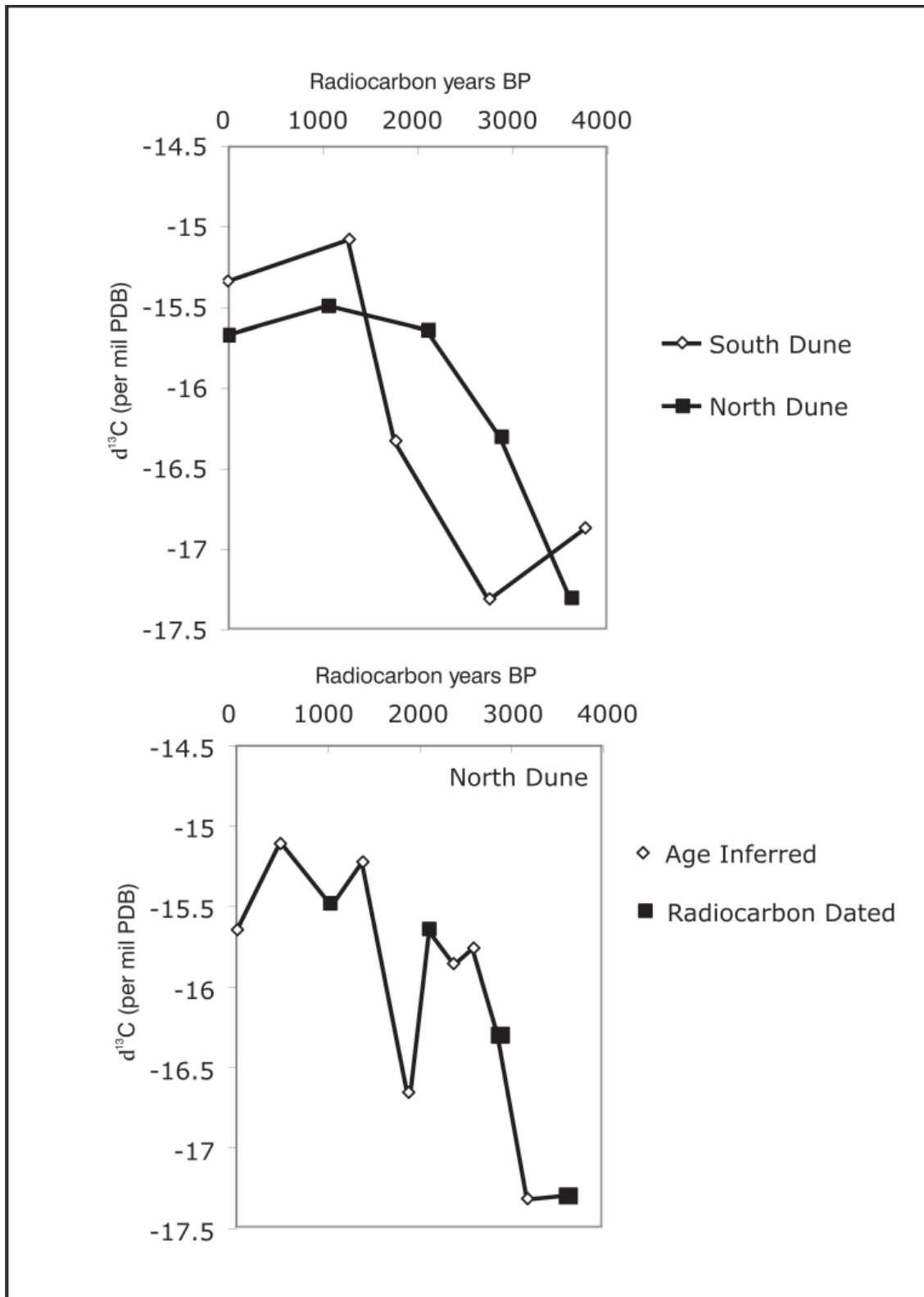


Figure F-5. Upper panel. Plot of $\delta^{13}C$ for radiocarbon dated soils in the north dune and the south dune at 41NU2. The most recent values were both the uppermost soil in the clay dune sequence and were not radiocarbon dated but assumed to reflect modern values. Lower panel: Plot of the $\delta^{13}C$ for all of the soils in the North Dune profile with inferred ages used to plot the soils for which radiocarbon ages are not available.

a climatic proxy. Detailed study of the lithological record of two dunes at the Cayo del Oso Site which have formed from a single mudflat on the western side of the Blind Oso, reveals that this hypothesis must be rejected. This work suggests that the lithological record of coastal clay dunes is intimately tied to the immediately adjacent wind-tidal flay and that even adjacent dunes may have significantly different sedimentary records which are largely driven by local factors other than climate. Microfossils which may provide details of the local sediment source returned variable results, with diatoms rarely preserved, and calcareous microfossils such as ostracodes and foraminifera

best preserved in soil C horizons. Paleosol A horizons were typically leached of calcium carbonate and rarely preserved significant numbers of calcareous microfossils.

The most climatically sensitive record present within these deposits is the record of stable carbon isotopic composition of the paleosol organic matter. The soil carbon isotopic records obtained from the two dunes chart similar trajectories throughout the last 4,000 years, and appear to provide a fairly sensitive (yet basic) record of local vegetation change on the dunes.

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APPENDIX G:
Radiocarbon Results

Darden Hood

FROM: Darden Hood, Director (mailto:<mailto:dhood@radiocarbon.com>)
(This is a copy of the letter being mailed. Invoices/receipts follow only by mail.)

November 13, 2007

Dr. James Abbott
Texas Department of Transportation
Cultural Resource Management
Environmental Affairs Division
125 East 11th Street
Austin, TX 78701

RE: Radiocarbon Dating Results For Samples 41NU2-067, 41NU2-069, 41NU2-071, 41NU2-072, 41NU2-153, 41NU2-156, 41NU2-157, 41NU2-158, 41NU2-259, 41NU2-260, 41NU2-261, 41NU2-262, 41NU2-371, 41NU2-373, 41NU2-375, 41NU2-376, 41NU2-377, 41NU2-380, 41NU2-381, 41NU2-382, 41NU2-384, 41NU2-387, 41NU2-388, 41NU2-389, 41NU2-391, 41NU2-393, 41NU2-396, 41NU2-397, 41NU2-398, 41NU2-399, 41NU2-400, 41NU2-403, 41NU2-405, 41NU2-412, 41NU2-415, 41NU2-416, 41NU2-417, 41NU2-418, 41NU2-419, 41NU2-420, 41NU2-422, 41NU2-423, 41NU2-424, 41NU2-425, 41NU2-426, 41NU2-427, 41NU2-428, 41NU2-429, 41NU2-430, 41NU2-431, 41NU2-432, 41NU2-433, 41NU2-434, 41NU2-436, 41NU2-437, 41NU2-438, 41NU2-439, 41NU2-440

Dear Jim:

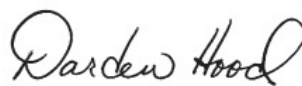
Enclosed are the radiocarbon dating results for 58 samples recently sent to us. They each provided plenty of carbon for accurate measurements and all the analyses proceeded normally. As usual, the method of analysis is listed on the report with the results and calibration data is provided where applicable.

As always, no students or intern researchers who would necessarily be distracted with other obligations and priorities were used in the analyses. We analyzed them with the combined attention of our entire professional staff.

If you have specific questions about the analyses, please contact us. We are always available to answer your questions.

Our invoice is enclosed. Please, forward it to the appropriate officer or send VISA charge authorization. Thank you. As always, if you have any questions or would like to discuss the results, don't hesitate to contact me.

Sincerely,



Dr. James Abbott

Report Date: 11/13/2007

Texas Department of Transportation

Material Received: 10/11/2007

Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235799 SAMPLE : 41NU2-067 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal AD 1310 to 1360 (Cal BP 640 to 590) AND Cal AD 1380 to 1440 (Cal BP 570 to 510)	400 +/- 40 BP	-16.1 o/oo	550 +/- 40 BP
Beta - 235800 SAMPLE : 41NU2-069 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal AD 780 to 1000 (Cal BP 1160 to 950)	970 +/- 40 BP	-15.4 o/oo	1130 +/- 40 BP
Beta - 235801 SAMPLE : 41NU2-071 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 390 to 200 (Cal BP 2340 to 2150)	2090 +/- 40 BP	-15.6 o/oo	2240 +/- 40 BP
Beta - 235802 SAMPLE : 41NU2-072 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 1360 to 1350 (Cal BP 3310 to 3300) AND Cal BC 1310 to 1050 (Cal BP 3260 to 3000)	2830 +/- 40 BP	-16.6 o/oo	2970 +/- 40 BP
Beta - 235803 SAMPLE : 41NU2-153 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal AD 890 to 1030 (Cal BP 1060 to 920)	900 +/- 40 BP	-15.5 o/oo	1060 +/- 40 BP

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Report Date: 11/13/2007

Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235804 SAMPLE : 41NU2-156 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 340 to 320 (Cal BP 2290 to 2270) AND Cal BC 210 to 40 (Cal BP 2160 to 1990)	2030 +/- 40 BP	-20.2 o/oo	2110 +/- 40 BP
Beta - 235805 SAMPLE : 41NU2-157 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 1200 to 930 (Cal BP 3150 to 2880)	2710 +/- 40 BP	-14.8 o/oo	2880 +/- 40 BP
Beta - 235806 SAMPLE : 41NU2-158 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 2130 to 2090 (Cal BP 4080 to 4040) AND Cal BC 2050 to 1890 (Cal BP 4000 to 3840)	3470 +/- 40 BP	-16.0 o/oo	3620 +/- 40 BP
Beta - 235807 SAMPLE : 41NU2-259 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal AD 660 to 810 (Cal BP 1290 to 1140)	1140 +/- 40 BP	-16.2 o/oo	1280 +/- 40 BP
Beta - 235808 SAMPLE : 41NU2-260 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal AD 140 to 380 (Cal BP 1810 to 1570)	1650 +/- 40 BP	-17.5 o/oo	1770 +/- 40 BP

Dr. James Abbott

Report Date: 11/13/2007

Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235809 SAMPLE : 41NU2-261 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 1000 to 820 (Cal BP 2950 to 2770)	2690 +/- 40 BP	-20.8 o/oo	2760 +/- 40 BP
Beta - 235810 SAMPLE : 41NU2-262 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (humate): acid washes 2 SIGMA CALIBRATION : Cal BC 2300 to 2120 (Cal BP 4250 to 4070) AND Cal BC 2090 to 2040 (Cal BP 4040 to 3990)	3650 +/- 40 BP	-17.7 o/oo	3770 +/- 40 BP
Beta - 235811 SAMPLE : 41NU2-371 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal AD 780 to 990 (Cal BP 1170 to 960)	1130 +/- 40 BP	-24.1 o/oo	1140 +/- 40 BP
Beta - 235812 SAMPLE : 41NU2-373 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal BC 410 to 360 (Cal BP 2360 to 2310) AND Cal BC 280 to 260 (Cal BP 2230 to 2200)	2290 +/- 40 BP	-23.5 o/oo	2310 +/- 40 BP
Beta - 235813 SAMPLE : 41NU2-375 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 650 to 770 (Cal BP 1300 to 1180)	1290 +/- 40 BP	-23.0 o/oo	1320 +/- 40 BP

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Report Date: 11/13/2007

Sample Data	Measured Radiocarbon Age	¹³ C/ ¹² C Ratio	Conventional Radiocarbon Age(*)
Beta - 235814 SAMPLE : 41NU2-376 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 10 to 210 (Cal BP 1940 to 1740)	1760 +/- 40 BP	-16.0 o/oo	1910 +/- 40 BP
Beta - 235815 SAMPLE : 41NU2-377 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 130 to 350 (Cal BP 1820 to 1600)	1670 +/- 40 BP	-18.4 o/oo	1780 +/- 40 BP
Beta - 235816 SAMPLE : 41NU2-380 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 100 to Cal AD 70 (Cal BP 2050 to 1880)	1990 +/- 40 BP	-23.6 o/oo	2010 +/- 40 BP
Beta - 235817 SAMPLE : 41NU2-381 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 80 to 250 (Cal BP 1870 to 1700)	1780 +/- 40 BP	-21.3 o/oo	1840 +/- 40 BP
Beta - 235818 SAMPLE : 41NU2-382 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 660 to 780 (Cal BP 1290 to 1160)	1230 +/- 40 BP	-21.1 o/oo	1290 +/- 40 BP

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Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235820 SAMPLE : 41NU2-384 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal BC 1120 to 890 (Cal BP 3070 to 2840) AND Cal BC 870 to 850 (Cal BP 2820 to 2800)	2780 +/- 50 BP	-22.1 o/oo	2830 +/- 50 BP
Beta - 235821 SAMPLE : 41NU2-387 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 790 to 510 (Cal BP 2740 to 2460)	2520 +/- 40 BP	-25.8 o/oo	2510 +/- 40 BP
Beta - 235822 SAMPLE : 41NU2-388 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 20 to 220 (Cal BP 1930 to 1730)	1840 +/- 40 BP	-21.1 o/oo	1900 +/- 40 BP
Beta - 235823 SAMPLE : 41NU2-389 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 30 to Cal AD 130 (Cal BP 1980 to 1820)	1920 +/- 40 BP	-23.8 o/oo	1940 +/- 40 BP
Beta - 235824 SAMPLE : 41NU2-391 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal BC 350 to 300 (Cal BP 2300 to 2260) AND Cal BC 210 to 40 (Cal BP 2160 to 1990)	2010 +/- 40 BP	-18.0 o/oo	2120 +/- 40 BP

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Sample Data	Measured Radiocarbon Age	¹³ C/ ¹² C Ratio	Conventional Radiocarbon Age(*)
Beta - 235825 SAMPLE : 41NU2-393 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal AD 540 to 650 (Cal BP 1410 to 1300)	1360 +/- 40 BP	-18.3 o/oo	1470 +/- 40 BP
Beta - 235826 SAMPLE : 41NU2-396 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 1680 to 1500 (Cal BP 3630 to 3440)	3280 +/- 40 BP	-23.7 o/oo	3300 +/- 40 BP
Beta - 235827 SAMPLE : 41NU2-397 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal BC 820 to 760 (Cal BP 2770 to 2720)	2510 +/- 40 BP	-19.1 o/oo	2610 +/- 40 BP
Beta - 235828 SAMPLE : 41NU2-398 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 780 to 410 (Cal BP 2730 to 2360)	2440 +/- 40 BP	-23.3 o/oo	2470 +/- 40 BP
Beta - 235829 SAMPLE : 41NU2-399 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 370 to 150 (Cal BP 2320 to 2100) AND Cal BC 140 to 110 (Cal BP 2090 to 2060)	2190 +/- 40 BP	-25.7 o/oo	2180 +/- 40 BP

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Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235830 SAMPLE : 41NU2-400 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal AD 780 to 990 (Cal BP 1170 to 960)	1120 +/- 40 BP	-23.9 o/oo	1140 +/- 40 BP
Beta - 235831 SAMPLE : 41NU2-403 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal BC 360 to 290 (Cal BP 2300 to 2240) AND Cal BC 240 to 50 (Cal BP 2180 to 2000)	2040 +/- 40 BP	-18.8 o/oo	2140 +/- 40 BP
Beta - 235832 SAMPLE : 41NU2-405 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (organic sediment): acid washes 2 SIGMA CALIBRATION : Cal BC 530 to 390 (Cal BP 2480 to 2340)	2280 +/- 40 BP	-19.3 o/oo	2370 +/- 40 BP
Beta - 235833 SAMPLE : 41NU2-412 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 180 to Cal AD 20 (Cal BP 2130 to 1930)	2020 +/- 40 BP	-22.6 o/oo	2060 +/- 40 BP
Beta - 235834 SAMPLE : 41NU2-415 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 780 to 400 (Cal BP 2730 to 2350)	2440 +/- 50 BP	-24.1 o/oo	2450 +/- 50 BP

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Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235835 SAMPLE : 41NU2-416 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 760 to 400 (Cal BP 2710 to 2350)	2420 +/- 40 BP	-23.9 o/oo	2440 +/- 40 BP
Beta - 235836 SAMPLE : 41NU2-417 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal AD 600 to 680 (Cal BP 1350 to 1270)	1360 +/- 40 BP	-23.4 o/oo	1390 +/- 40 BP
Beta - 235837 SAMPLE : 41NU2-418 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (charred material): acid/alkali/acid 2 SIGMA CALIBRATION : Cal BC 1520 to 1400 (Cal BP 3470 to 3350)	3190 +/- 40 BP	-25.0 o/oo	3190 +/- 40 BP
Beta - 235838 SAMPLE : 41NU2-419 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 940 to 1070 (Cal BP 1010 to 880)	990 +/- 40 BP	-0.4 o/oo O18/O16 = -1.0 o/oo	1390 +/- 40 BP
Beta - 235839 SAMPLE : 41NU2-420 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 180 to 390 (Cal BP 1770 to 1560)	1690 +/- 40 BP	-0.4 o/oo O18/O16 = +0.1 o/oo	2090 +/- 40 BP

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Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235841	1680 +/- 40 BP	+0.8 o/oo O18/O16 = -0.9 o/oo	2100 +/- 40 BP
SAMPLE : 41NU2-422 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 160 to 370 (Cal BP 1790 to 1580)			
Beta - 235842	2130 +/- 40 BP	-0.2 o/oo O18/O16 = +0.5 o/oo	2540 +/- 40 BP
SAMPLE : 41NU2-423 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 360 to 160 (Cal BP 2310 to 2110)			
Beta - 235843	1780 +/- 40 BP	+0.5 o/oo O18/O16 = -1.1 o/oo	2200 +/- 40 BP
SAMPLE : 41NU2-424 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 70 to 250 (Cal BP 1880 to 1700)			
Beta - 235844	2450 +/- 40 BP	-0.6 o/oo O18/O16 = -1.1 o/oo	2850 +/- 40 BP
SAMPLE : 41NU2-425 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 770 to 530 (Cal BP 2720 to 2480)			
Beta - 235845	5130 +/- 40 BP	-0.3 o/oo O18/O16 = -1.6 o/oo	5540 +/- 40 BP
SAMPLE : 41NU2-426 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 4040 to 3920 (Cal BP 5990 to 5870)			

Dr. James Abbott

Report Date: 11/13/2007

Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235846	1700 +/- 40 BP	+0.2 o/oo O18/O16 = -0.8 o/oo	2110 +/- 40 BP
SAMPLE : 41NU2-427 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 150 to 360 (Cal BP 1800 to 1590)			
Beta - 235847	2160 +/- 40 BP	+1.0 o/oo O18/O16 = -0.2 o/oo	2590 +/- 40 BP
SAMPLE : 41NU2-428 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 390 to 200 (Cal BP 2340 to 2150)			
Beta - 235848	2690 +/- 40 BP	-0.4 o/oo O18/O16 = -1.0 o/oo	3090 +/- 40 BP
SAMPLE : 41NU2-429 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 1000 to 810 (Cal BP 2950 to 2760)			
Beta - 235849	1620 +/- 40 BP	+0.1 o/oo O18/O16 = +0.2 o/oo	2030 +/- 40 BP
SAMPLE : 41NU2-430 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 260 to 440 (Cal BP 1690 to 1510)			
Beta - 235850	1850 +/- 40 BP	+0.6 o/oo O18/O16 = -0.7 o/oo	2270 +/- 40 BP
SAMPLE : 41NU2-431 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 20 to Cal AD 150 (Cal BP 1970 to 1800)			

Dr. James Abbott

Report Date: 11/13/2007

Sample Data	Measured Radiocarbon Age	¹³ C/ ¹² C Ratio	Conventional Radiocarbon Age(*)
Beta - 235851	1030 +/- 40 BP	-0.5 o/oo O18/O16 = -1.0 o/oo	1430 +/- 40 BP
SAMPLE : 41NU2-432 ANALYSIS : Radiometric-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 900 to 1040 (Cal BP 1050 to 900)			
Beta - 235852	1910 +/- 40 BP	-0.8 o/oo O18/O16 = -0.9 o/oo	2310 +/- 40 BP
SAMPLE : 41NU2-433 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 60 to Cal AD 120 (Cal BP 2010 to 1840)			
Beta - 235853	2280 +/- 40 BP	+0.1 o/oo O18/O16 = +2.0 o/oo	2700 +/- 50 BP
SAMPLE : 41NU2-434 ANALYSIS : Radiometric-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 590 to 350 (Cal BP 2540 to 2300)			
Beta - 235855	940 +/- 40 BP	+1.1 o/oo O18/O16 = -0.6 o/oo	1370 +/- 40 BP
SAMPLE : 41NU2-436 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 970 to 1110 (Cal BP 980 to 840)			
Beta - 235856	1100 +/- 40 BP	-0.3 o/oo O18/O16 = 0.0 o/oo	1510 +/- 40 BP
SAMPLE : 41NU2-437 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal AD 790 to 990 (Cal BP 1160 to 960)			

Dr. James Abbott

Report Date: 11/13/2007

Sample Data	Measured Radiocarbon Age	$^{13}\text{C}/^{12}\text{C}$ Ratio	Conventional Radiocarbon Age(*)
Beta - 235857 SAMPLE : 41NU2-438 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 340 to 110 (Cal BP 2290 to 2060)	2080 +/- 40 BP	+0.6 o/oo O18/O16 = +0.2 o/oo	2500 +/- 40 BP
Beta - 235858 SAMPLE : 41NU2-439 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 330 to 80 (Cal BP 2280 to 2030)	2070 +/- 40 BP	-0.2 o/oo O18/O16 = -0.2 o/oo	2480 +/- 40 BP
Beta - 235859 SAMPLE : 41NU2-440 ANALYSIS : AMS-Standard delivery MATERIAL/PRETREATMENT : (shell): acid etch 2 SIGMA CALIBRATION : Cal BC 1680 to 1490 (Cal BP 3630 to 3440)	3230 +/- 40 BP	-0.1 o/oo O18/O16 = +1.4 o/oo	3640 +/- 40 BP

CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-16.1:lab. mult=1)

Laboratory number: **Beta-235799**

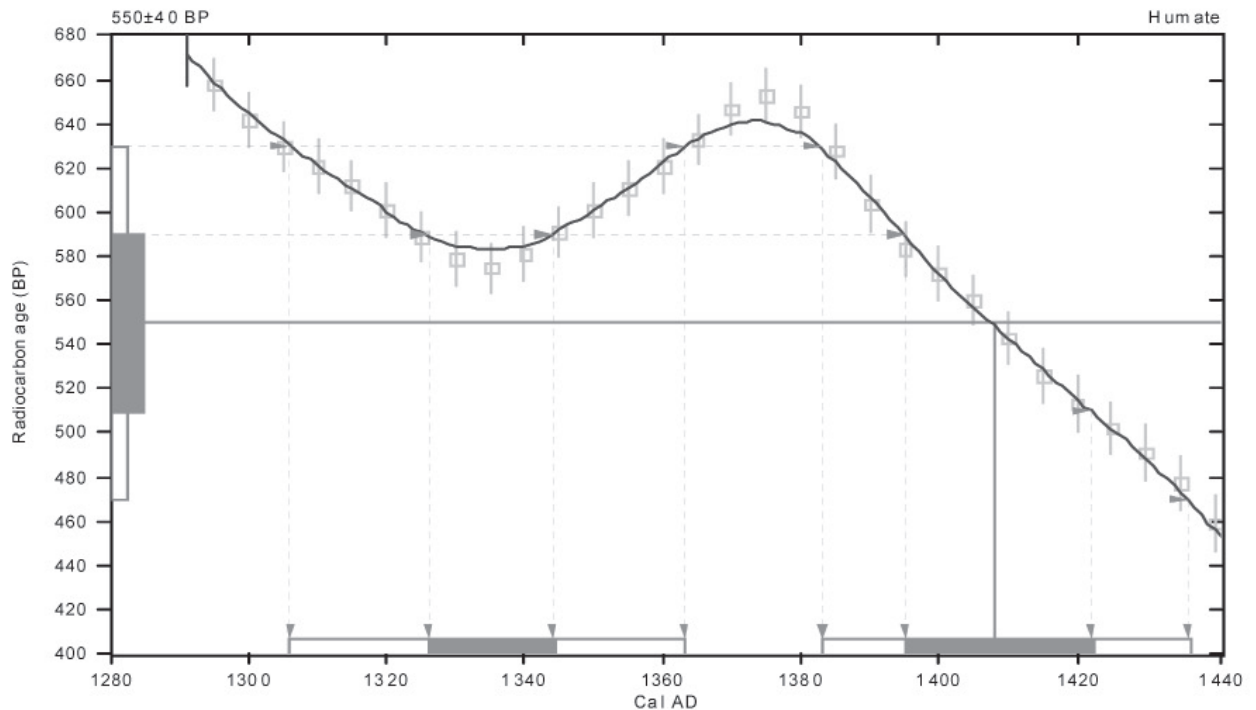
Conventional radiocarbon age: **550±40 BP**

2 Sigma calibrated results: Cal AD 1310 to 1360 (Cal BP 640 to 590) and
(95% probability) Cal AD 1380 to 1440 (Cal BP 570 to 510)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 1410 (Cal BP 540)

1 Sigma calibrated results: Cal AD 1330 to 1340 (Cal BP 620 to 610) and
(68% probability) Cal AD 1400 to 1420 (Cal BP 560 to 530)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-15.4:lab. mult=1)

Laboratory number: **Beta-235800**

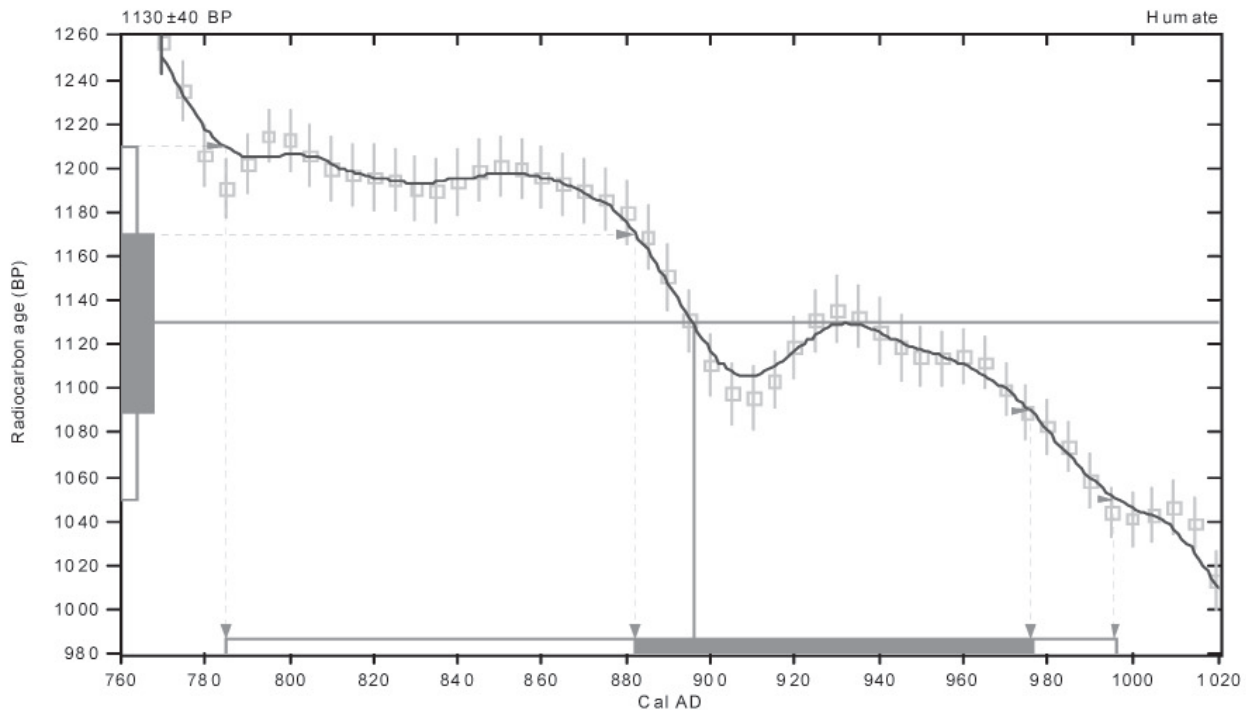
Conventional radiocarbon age: **1130±40 BP**

2 Sigma calibrated result: **Cal AD 780 to 1000 (Cal BP 1160 to 950)**
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 900 (Cal BP 1050)**

1 Sigma calibrated result: **Cal AD 880 to 980 (Cal BP 1070 to 970)**
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-15.6:lab. mult=1)

Laboratory number: **Beta-235801**

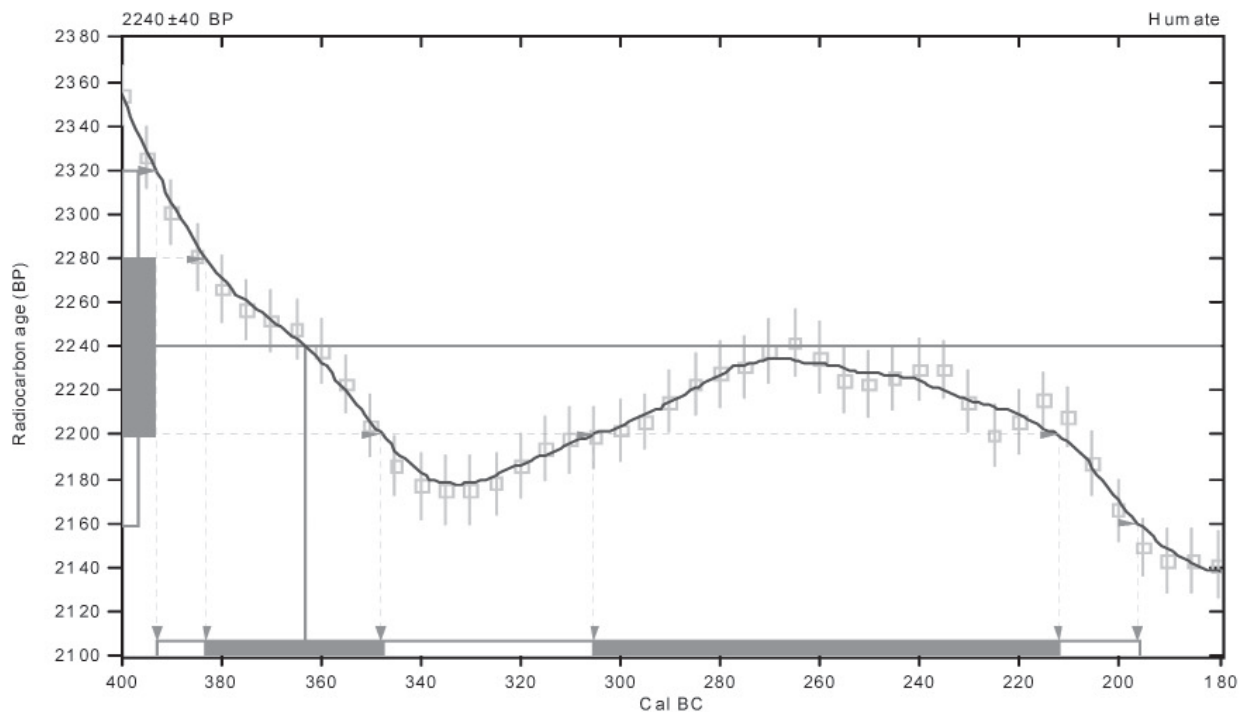
Conventional radiocarbon age: **2240±40 BP**

2 Sigma calibrated result: Cal BC 390 to 200 (Cal BP 2340 to 2150)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 360 (Cal BP 2310)

1 Sigma calibrated results: Cal BC 380 to 350 (Cal BP 2330 to 2300) and
(68% probability) **Cal BC 300 to 210 (Cal BP 2260 to 2160)**



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-16.6;lab. mult=1)

Laboratory number: **Beta-235802**

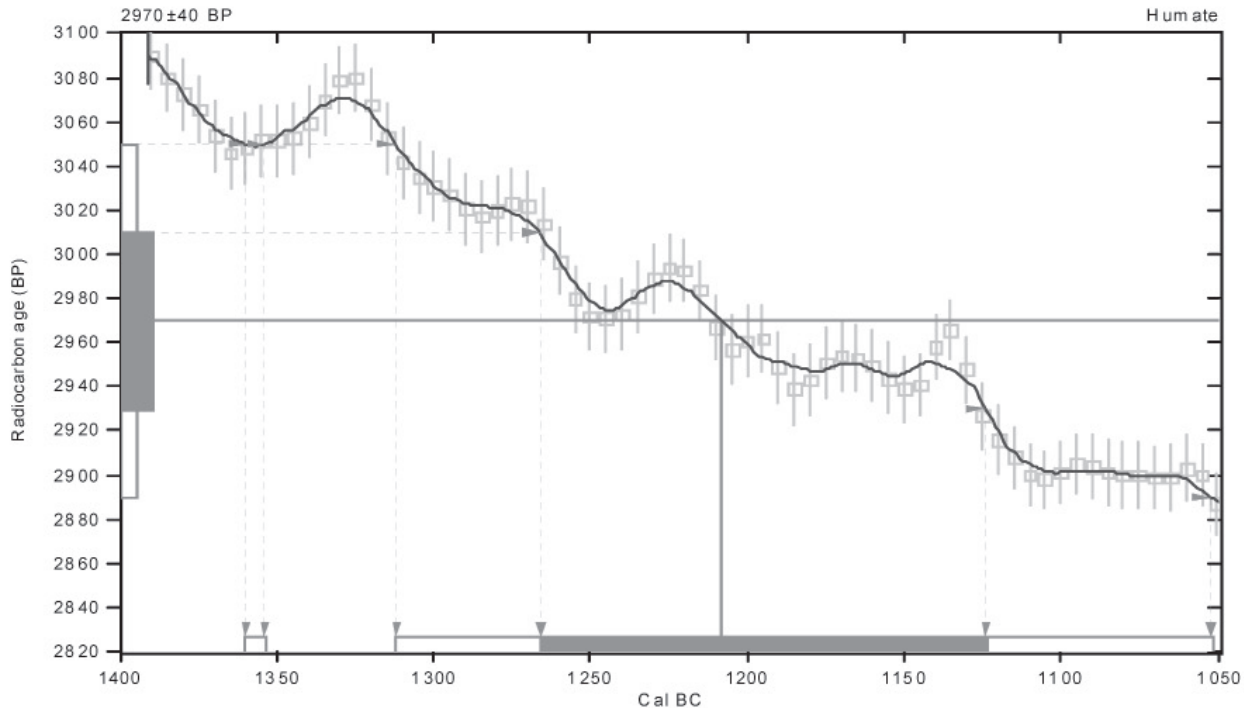
Conventional radiocarbon age: **2970±40 BP**

2 Sigma calibrated results: Cal BC 1360 to 1350 (Cal BP 3310 to 3300) and
(95% probability) Cal BC 1310 to 1050 (Cal BP 3260 to 3000)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 1210 (Cal BP 3160)

1 Sigma calibrated result: Cal BC 1270 to 1120 (Cal BP 3220 to 3070)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-15.5:lab. mult=1)

Laboratory number: **Beta-235803**

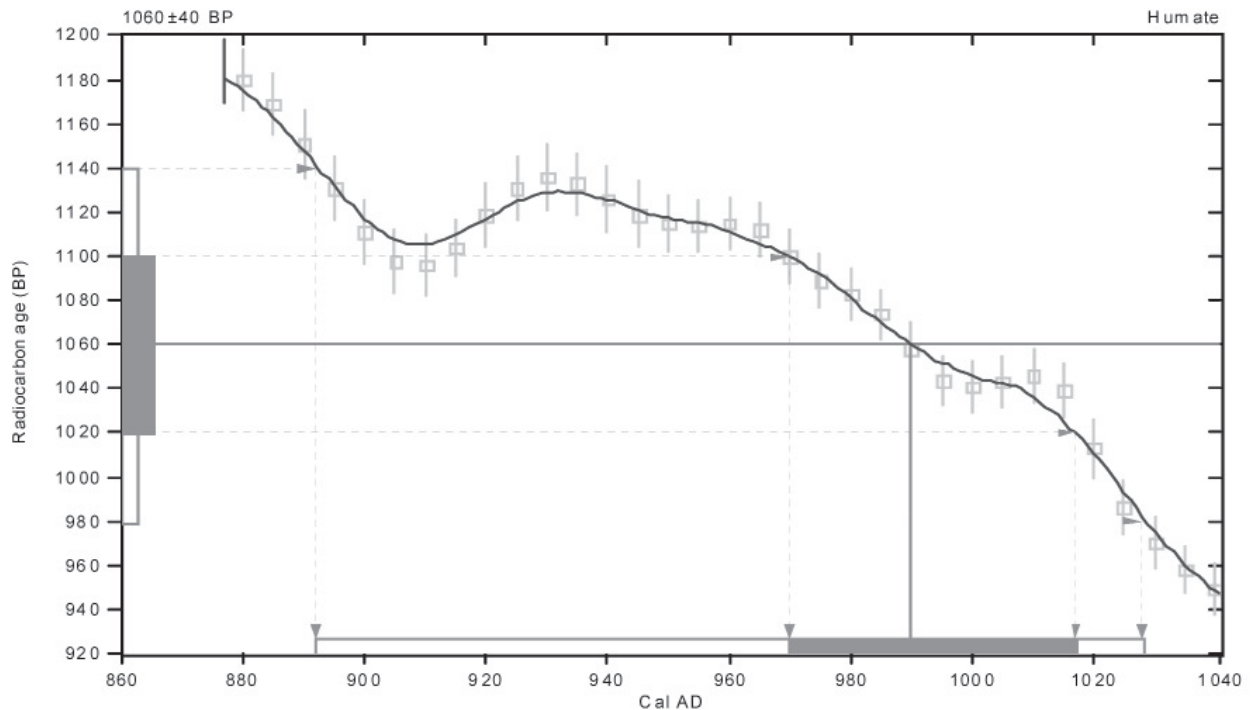
Conventional radiocarbon age: **1060±40 BP**

2 Sigma calibrated result: Cal AD 890 to 1030 (Cal BP 1060 to 920)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 990 (Cal BP 960)**

1 Sigma calibrated result: Cal AD 970 to 1020 (Cal BP 980 to 930)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-20.2:lab. mult=1)

Laboratory number: **Beta-235804**

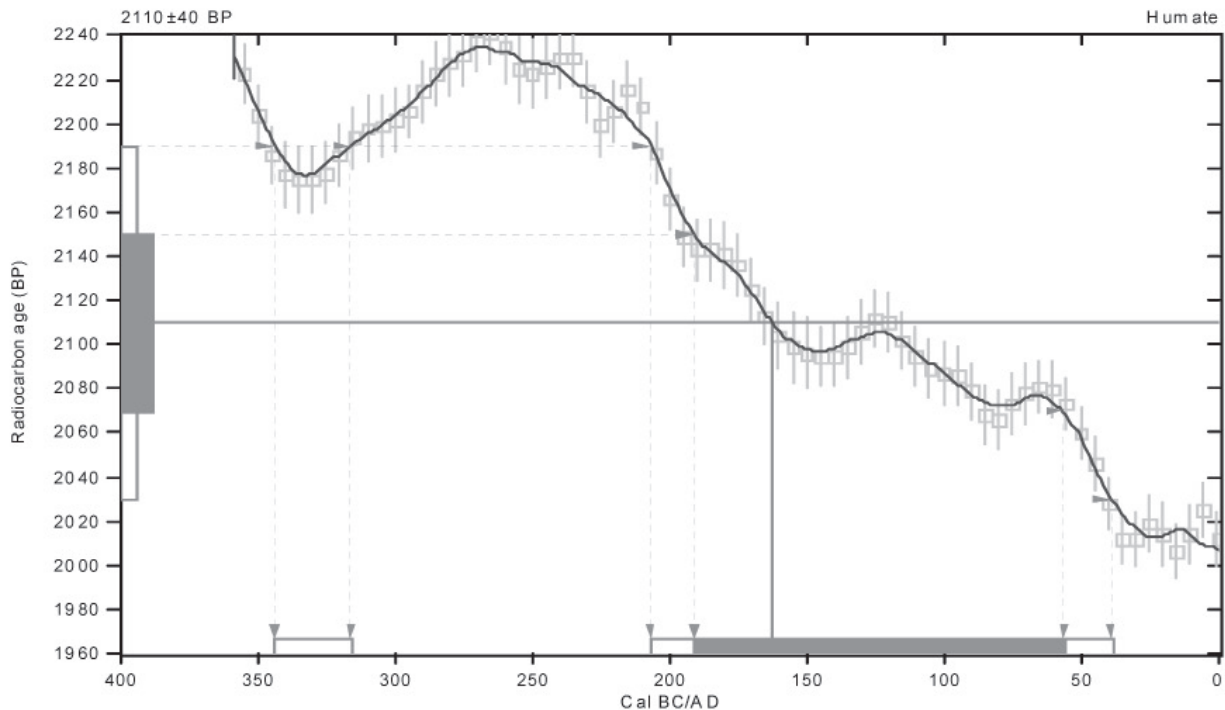
Conventional radiocarbon age: **2110±40 BP**

2 Sigma calibrated results: Cal BC 340 to 320 (Cal BP 2290 to 2270) and
(95% probability) Cal BC 210 to 40 (Cal BP 2160 to 1990)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 160 (Cal BP 2110)

1 Sigma calibrated result: Cal BC 190 to 60 (Cal BP 2140 to 2010)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-14.8:lab. mult=1)

Laboratory number: **Beta-235805**

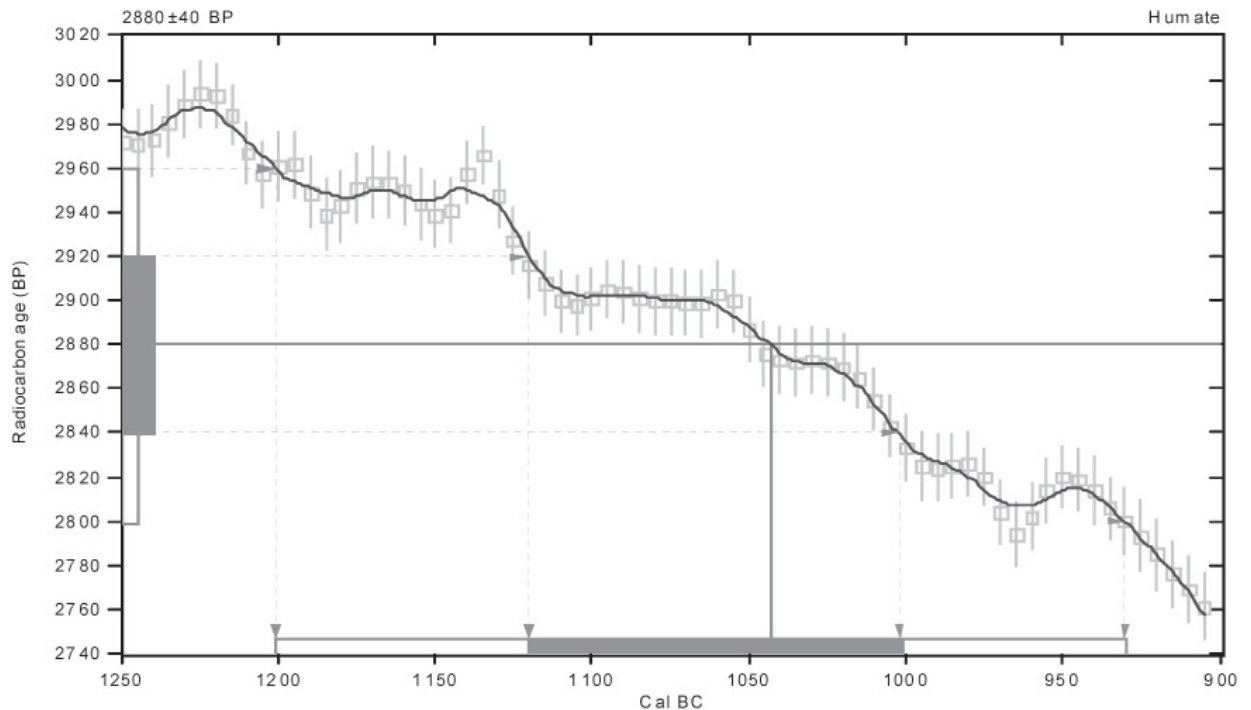
Conventional radiocarbon age: **2880±40 BP**

2 Sigma calibrated result: Cal BC 1200 to 930 (Cal BP 3150 to 2880)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 1040 (Cal BP 2990)**

1 Sigma calibrated result: Cal BC 1120 to 1000 (Cal BP 3070 to 2950)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-16;lab. mult=1)

Laboratory number: **Beta-235806**

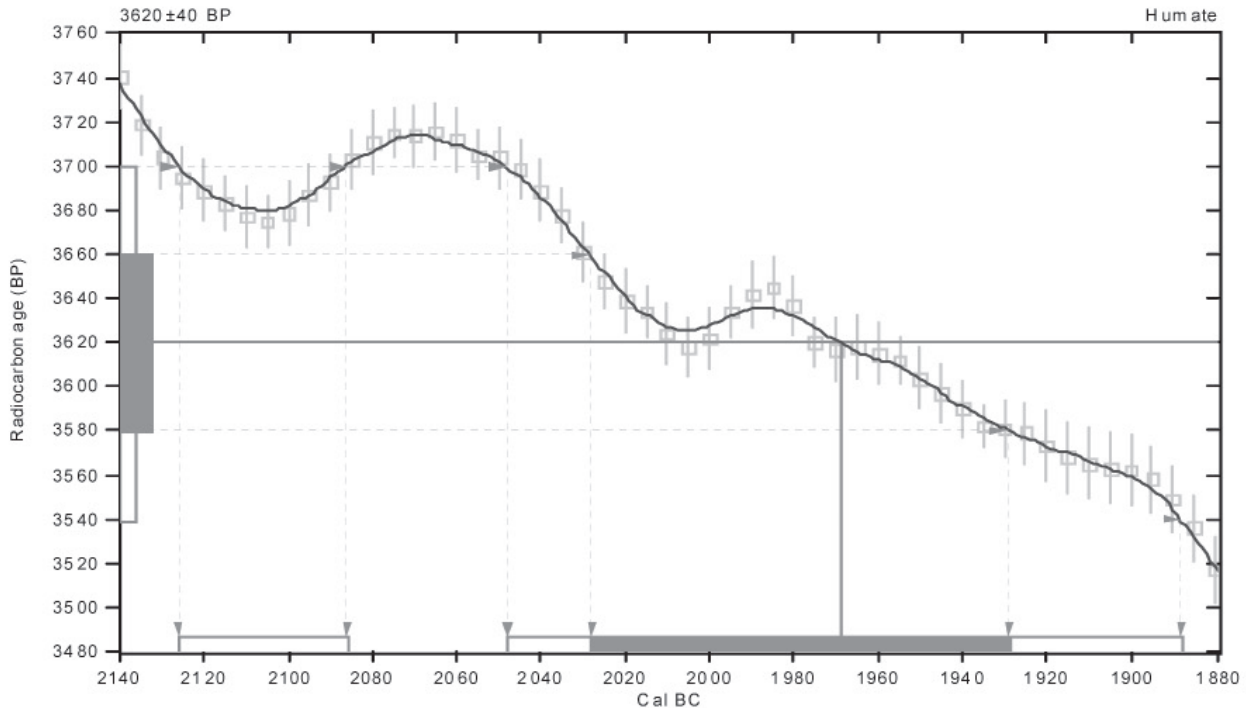
Conventional radiocarbon age: **3620±40 BP**

2 Sigma calibrated results: Cal BC 2130 to 2090 (Cal BP 4080 to 4040) and
(95% probability) Cal BC 2050 to 1890 (Cal BP 4000 to 3840)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 1970 (Cal BP 3920)

1 Sigma calibrated result: Cal BC 2030 to 1930 (Cal BP 3980 to 3880)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-16.2:lab. mult=1)

Laboratory number: **Beta-235807**

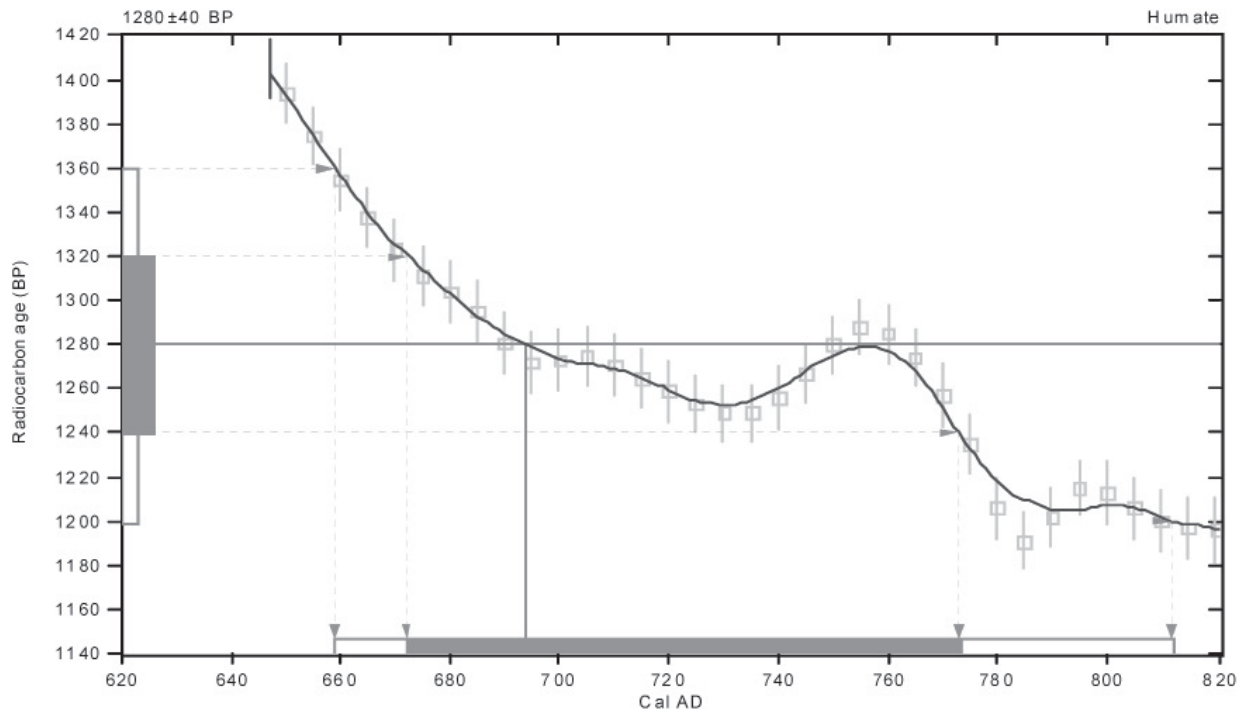
Conventional radiocarbon age: **1280±40 BP**

2 Sigma calibrated result: Cal AD 660 to 810 (Cal BP 1290 to 1140)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 690 (Cal BP 1260)

1 Sigma calibrated result: Cal AD 670 to 770 (Cal BP 1280 to 1180)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-17.5:lab. mult=1)

Laboratory number: **Beta-235808**

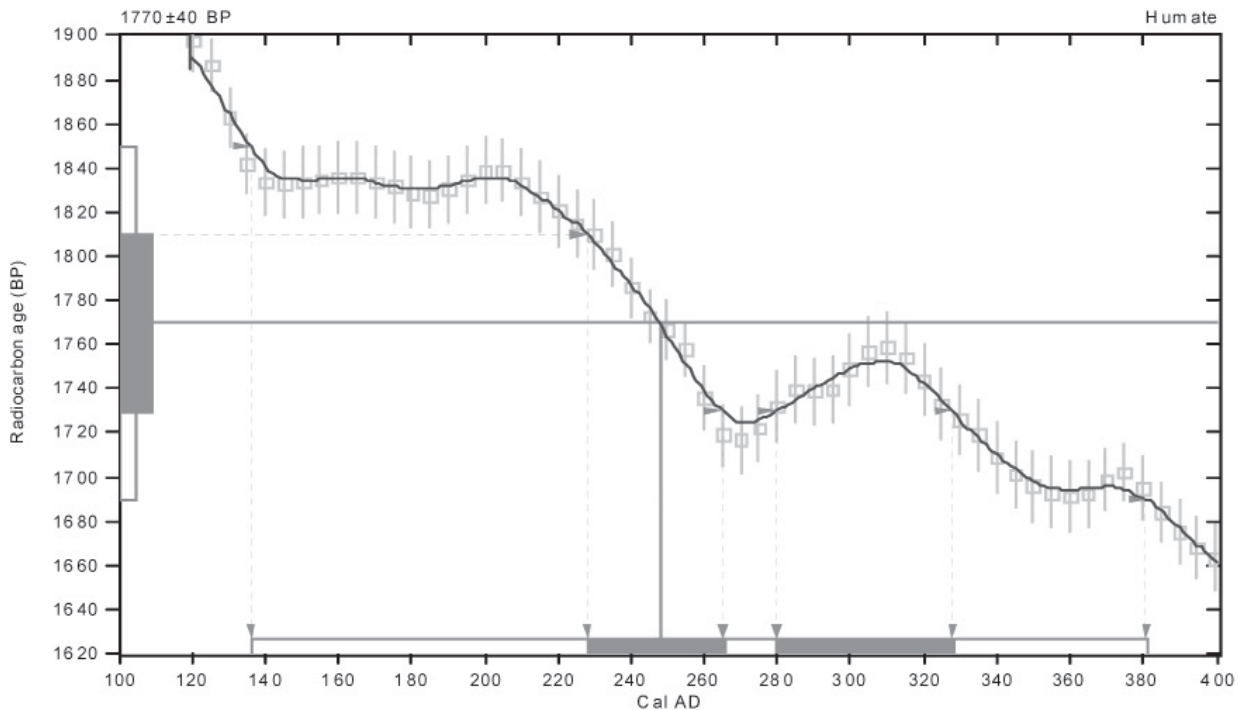
Conventional radiocarbon age: **1770±40 BP**

2 Sigma calibrated result: Cal AD 140 to 380 (Cal BP 1810 to 1570)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 250 (Cal BP 1700)

1 Sigma calibrated results: Cal AD 230 to 260 (Cal BP 1720 to 1680) and
(68% probability) **Cal AD 280 to 330 (Cal BP 1670 to 1620)**



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-20.8;lab. mult=1)

Laboratory number: **Beta-235809**

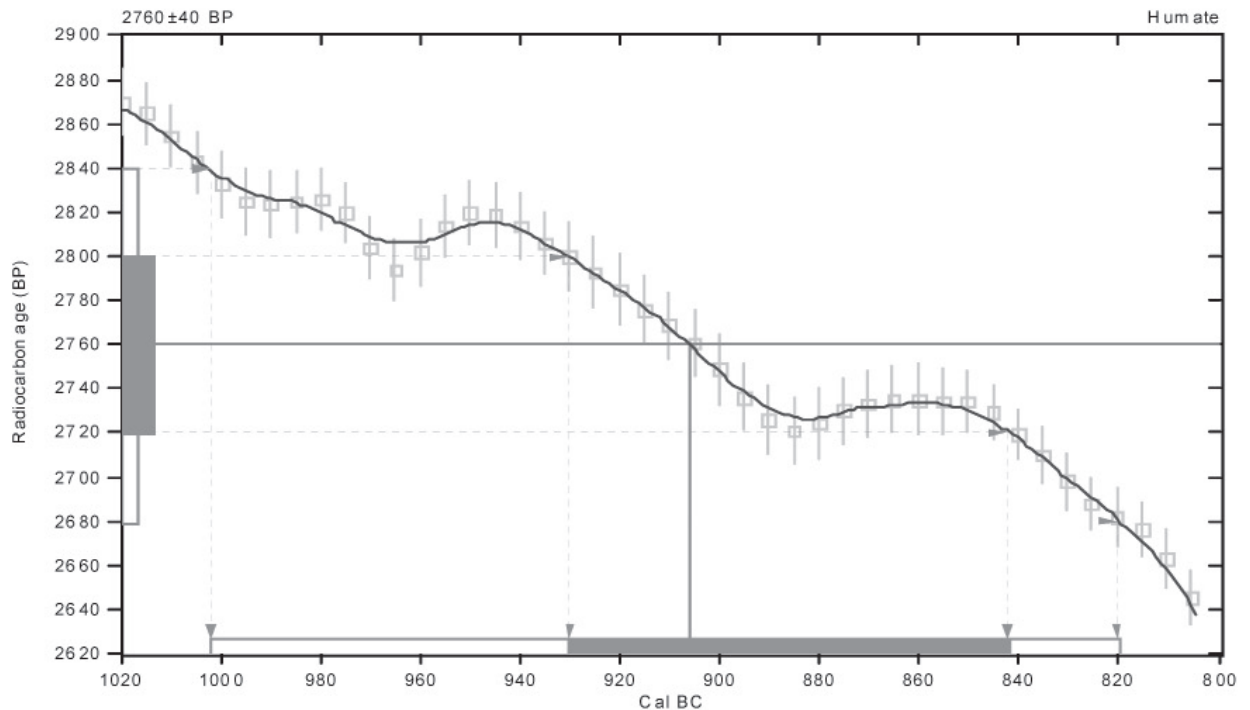
Conventional radiocarbon age: **2760±40 BP**

2 Sigma calibrated result: Cal BC 1000 to 820 (Cal BP 2950 to 2770)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 910 (Cal BP 2860)**

1 Sigma calibrated result: Cal BC 930 to 840 (Cal BP 2880 to 2790)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-17.7;lab. mult=1)

Laboratory number: **Beta-235810**

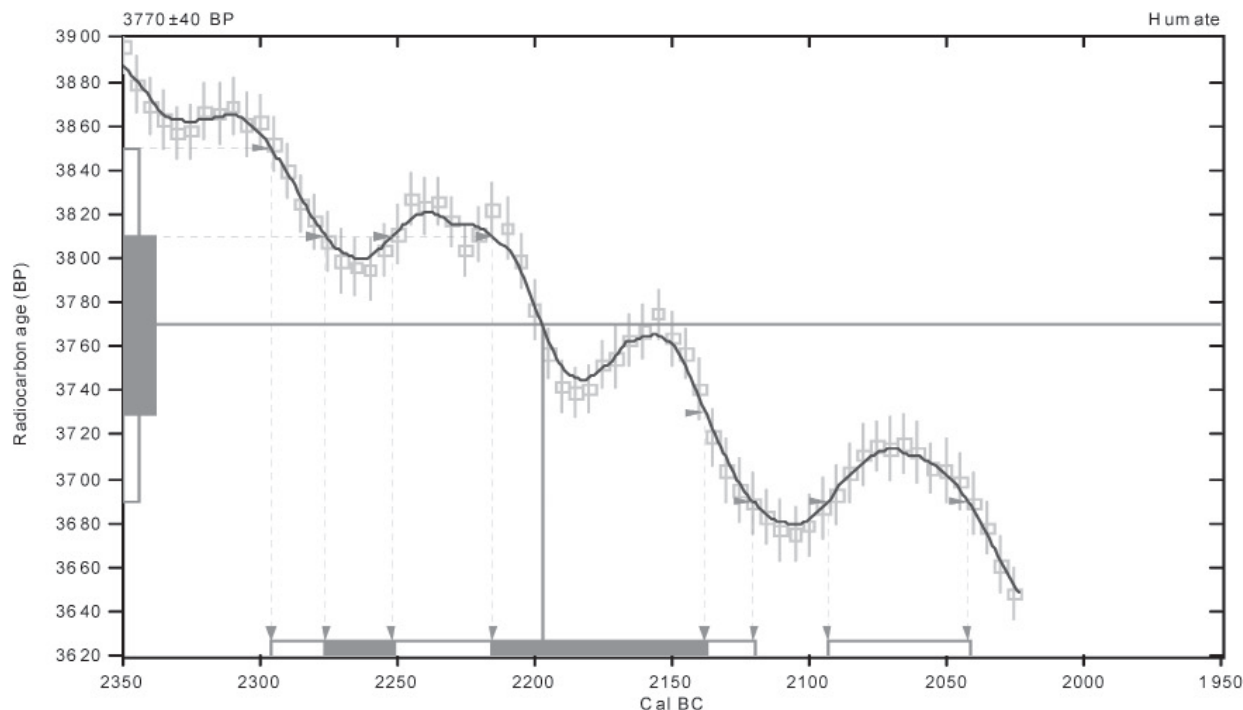
Conventional radiocarbon age: **3770±40 BP**

2 Sigma calibrated results: Cal BC 2300 to 2120 (Cal BP 4250 to 4070) and
(95% probability) Cal BC 2090 to 2040 (Cal BP 4040 to 3990)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 2200 (Cal BP 4150)

1 Sigma calibrated results: Cal BC 2280 to 2250 (Cal BP 4230 to 4200) and
(68% probability) Cal BC 2220 to 2140 (Cal BP 4160 to 4090)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-24.1:lab. mult=1)

Laboratory number: **Beta-235811**

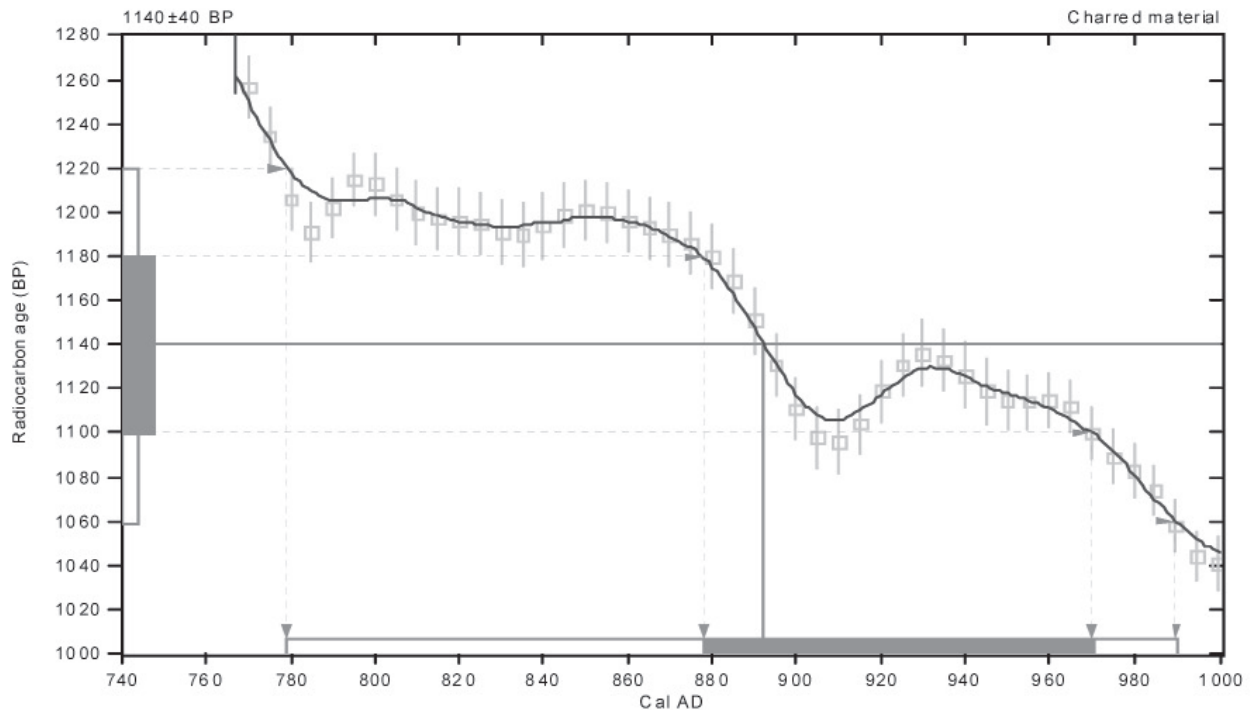
Conventional radiocarbon age: **1140±40 BP**

2 Sigma calibrated result: Cal AD 780 to 990 (Cal BP 1170 to 960)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 890 (Cal BP 1060)

1 Sigma calibrated result: Cal AD 880 to 970 (Cal BP 1070 to 980)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.5:lab. mult=1)

Laboratory number: **Beta-235812**

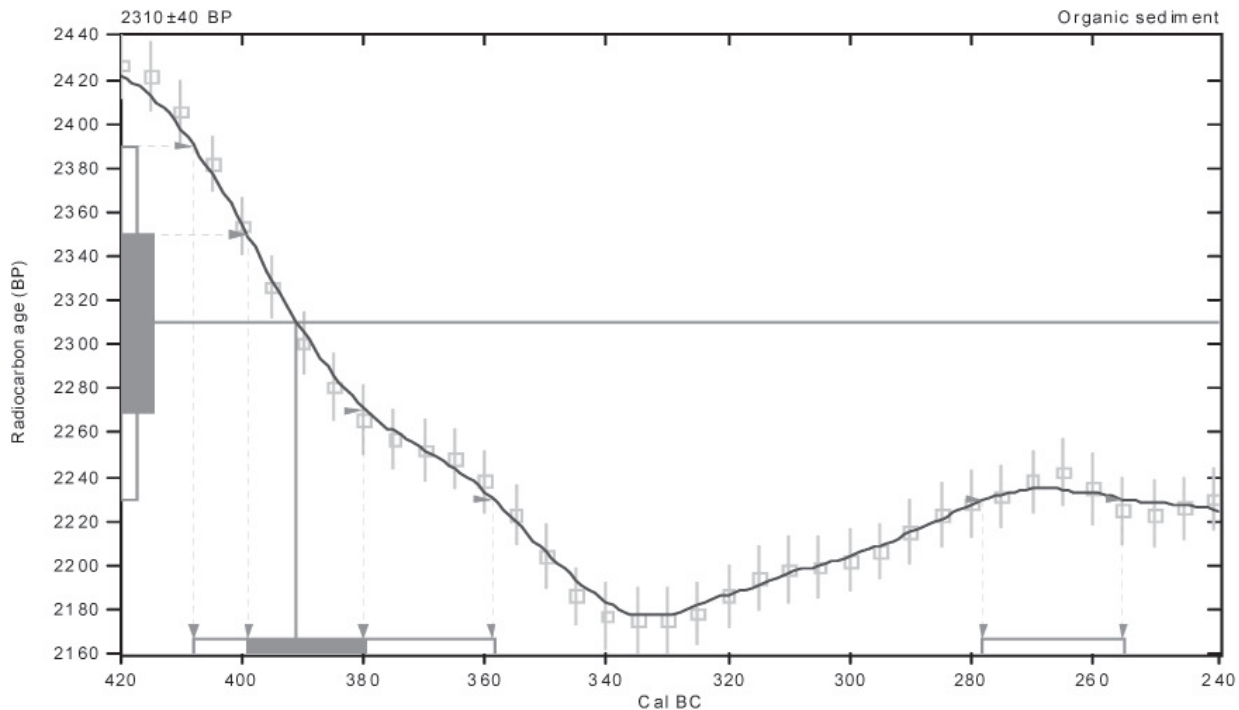
Conventional radiocarbon age: **2310±40 BP**

2 Sigma calibrated results: Cal BC 410 to 360 (Cal BP 2360 to 2310) and
(95% probability) Cal BC 280 to 260 (Cal BP 2230 to 2200)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 390 (Cal BP 2340)

1 Sigma calibrated result: Cal BC 400 to 380 (Cal BP 2350 to 2330)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23:lab. mult=1)

Laboratory number: **Beta-235813**

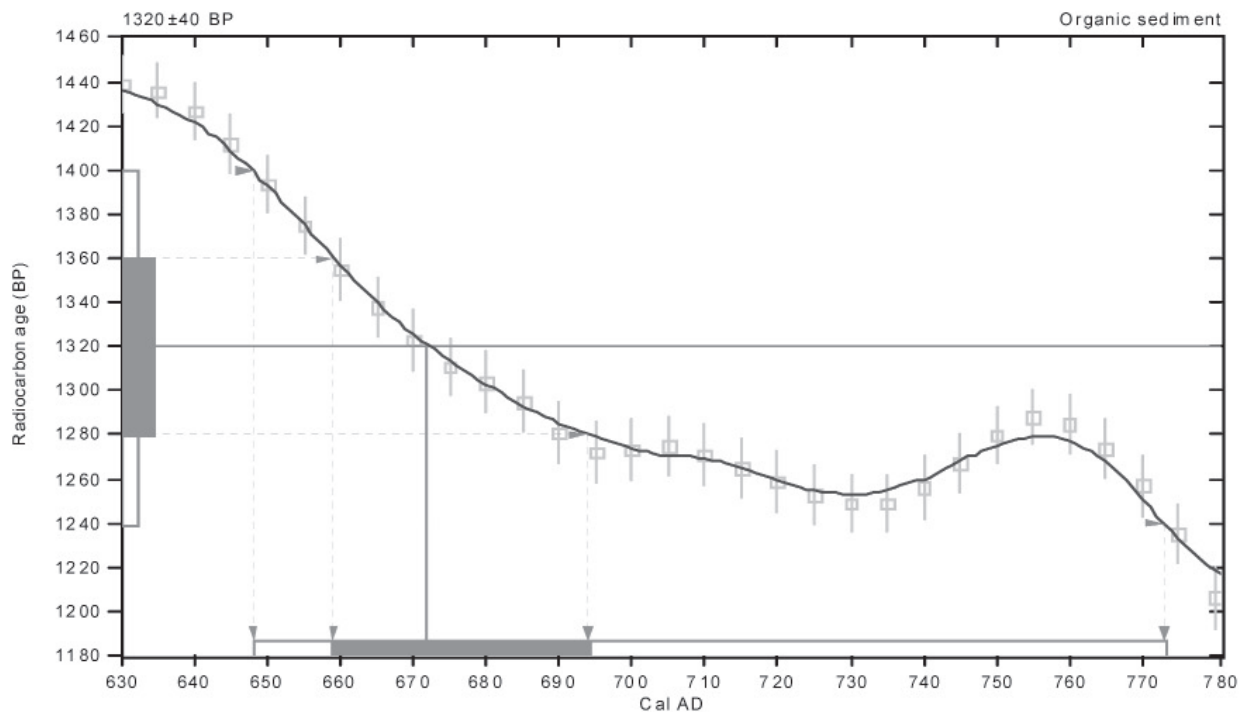
Conventional radiocarbon age: **1320±40 BP**

2 Sigma calibrated result: Cal AD 650 to 770 (Cal BP 1300 to 1180)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 670 (Cal BP 1280)

1 Sigma calibrated result: Cal AD 660 to 690 (Cal BP 1290 to 1260)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-16;lab. mult=1)

Laboratory number: Beta-235814

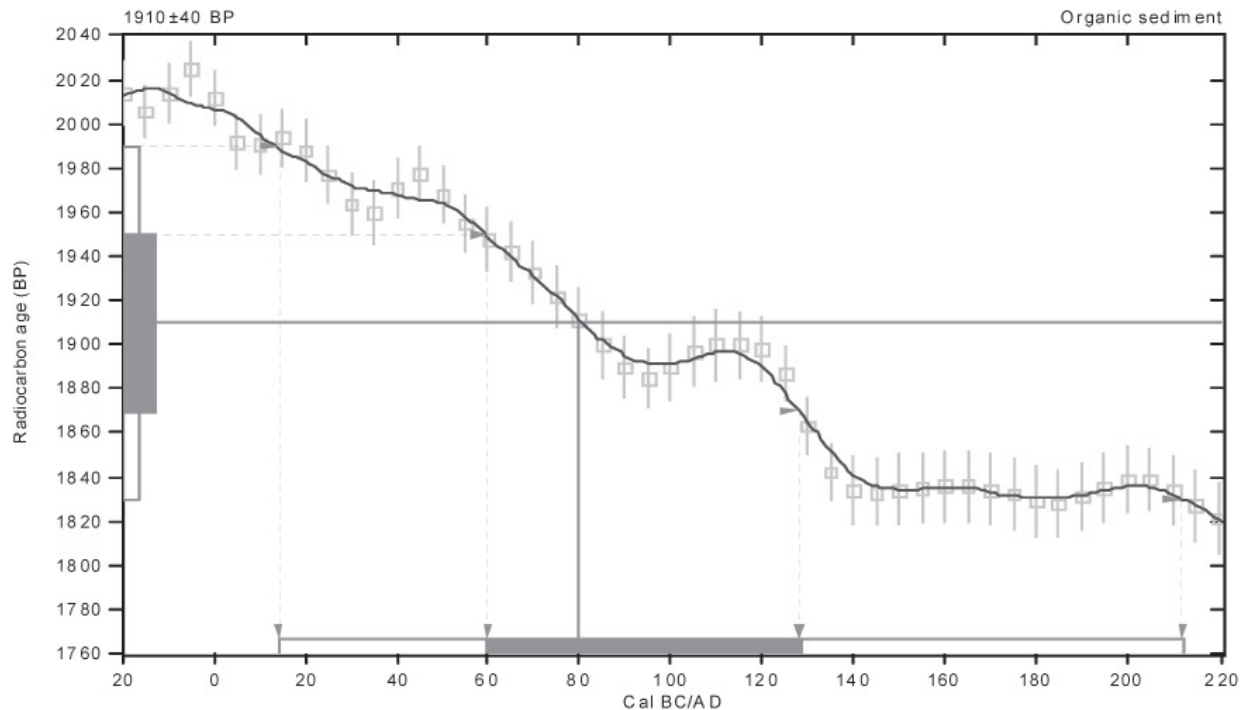
Conventional radiocarbon age: 1910±40 BP

2 Sigma calibrated result: Cal AD 10 to 210 (Cal BP 1940 to 1740)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 80 (Cal BP 1870)

1 Sigma calibrated result: Cal AD 60 to 130 (Cal BP 1890 to 1820)
(68% probability)



References:

Database used
INTCAL04

Calibration Database
INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-18.4:lab. mult=1)

Laboratory number: **Beta-235815**

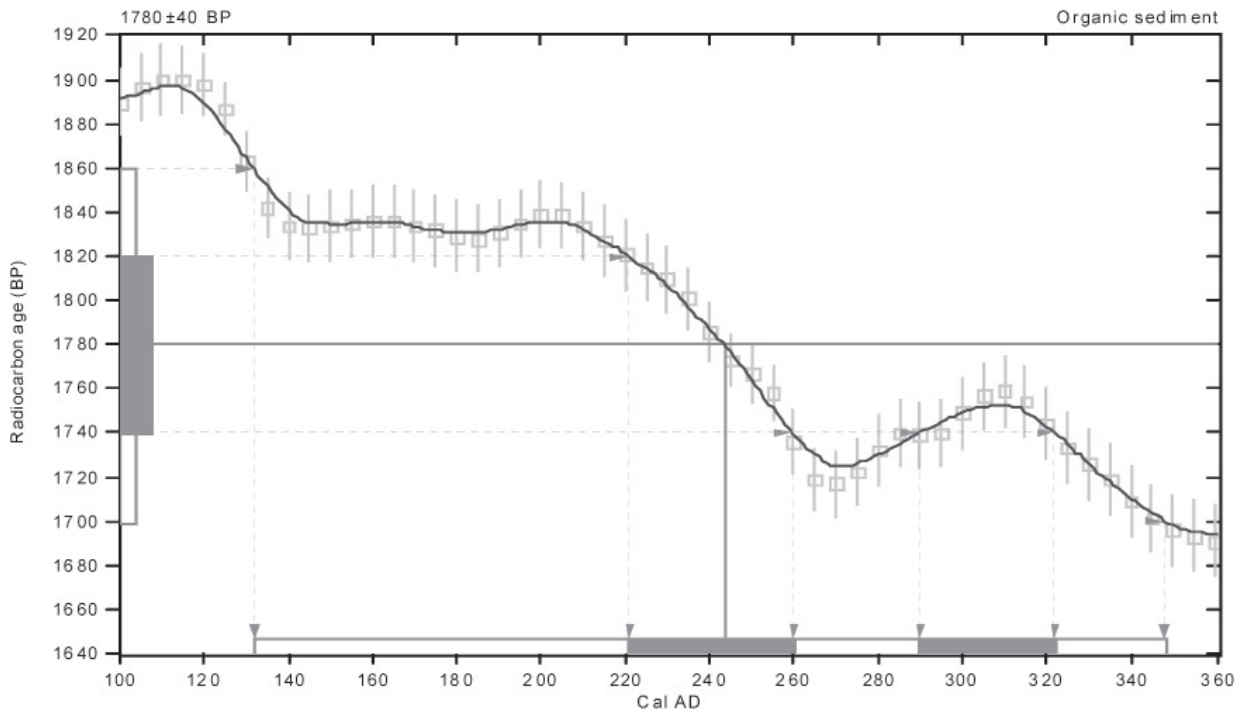
Conventional radiocarbon age: **1780±40 BP**

2 Sigma calibrated result: Cal AD 130 to 350 (Cal BP 1820 to 1600)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 240 (Cal BP 1710)

1 Sigma calibrated results: Cal AD 220 to 260 (Cal BP 1730 to 1690) and
(68% probability) **Cal AD 290 to 320 (Cal BP 1660 to 1630)**



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

Beta Analytic Radiocarbon Dating Laboratory

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.6:lab. mult=1)

Laboratory number: **Beta-235816**

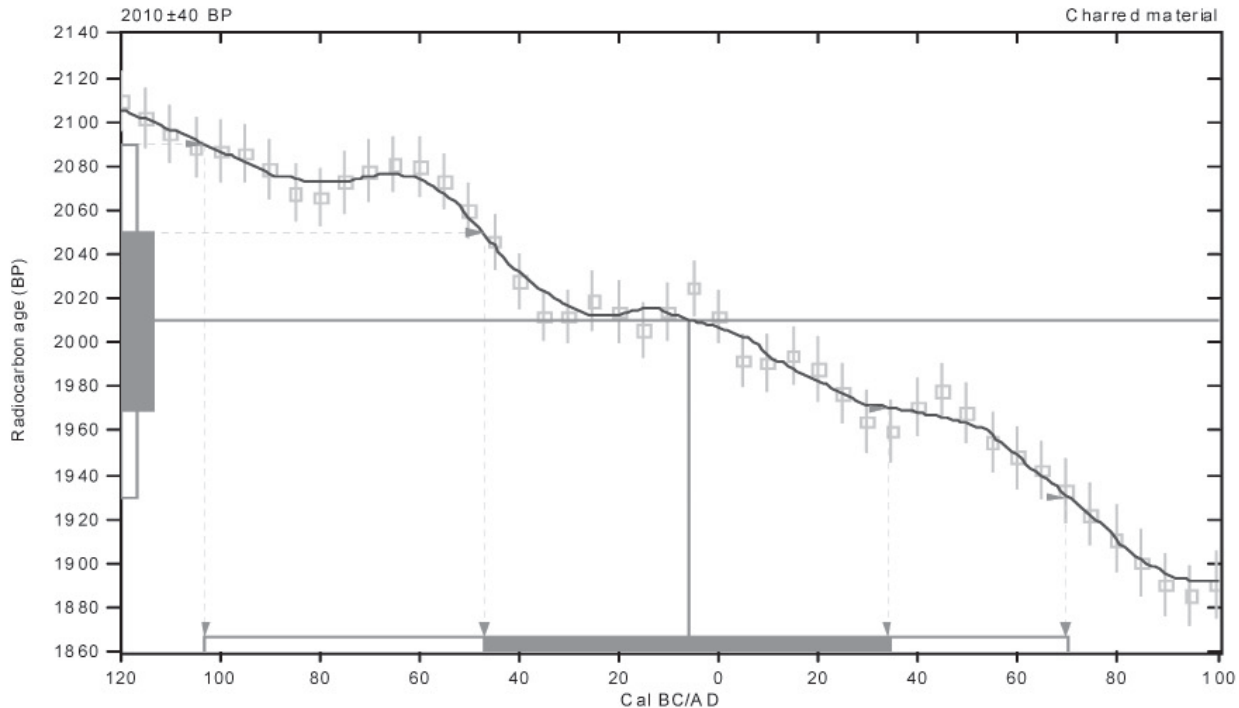
Conventional radiocarbon age: **2010±40 BP**

2 Sigma calibrated result: Cal BC 100 to Cal AD 70 (Cal BP 2050 to 1880)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 10 (Cal BP 1960)**

1 Sigma calibrated result: Cal BC 50 to Cal AD 30 (Cal BP 2000 to 1920)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-21.3:lab. mult=1)

Laboratory number: Beta-235817

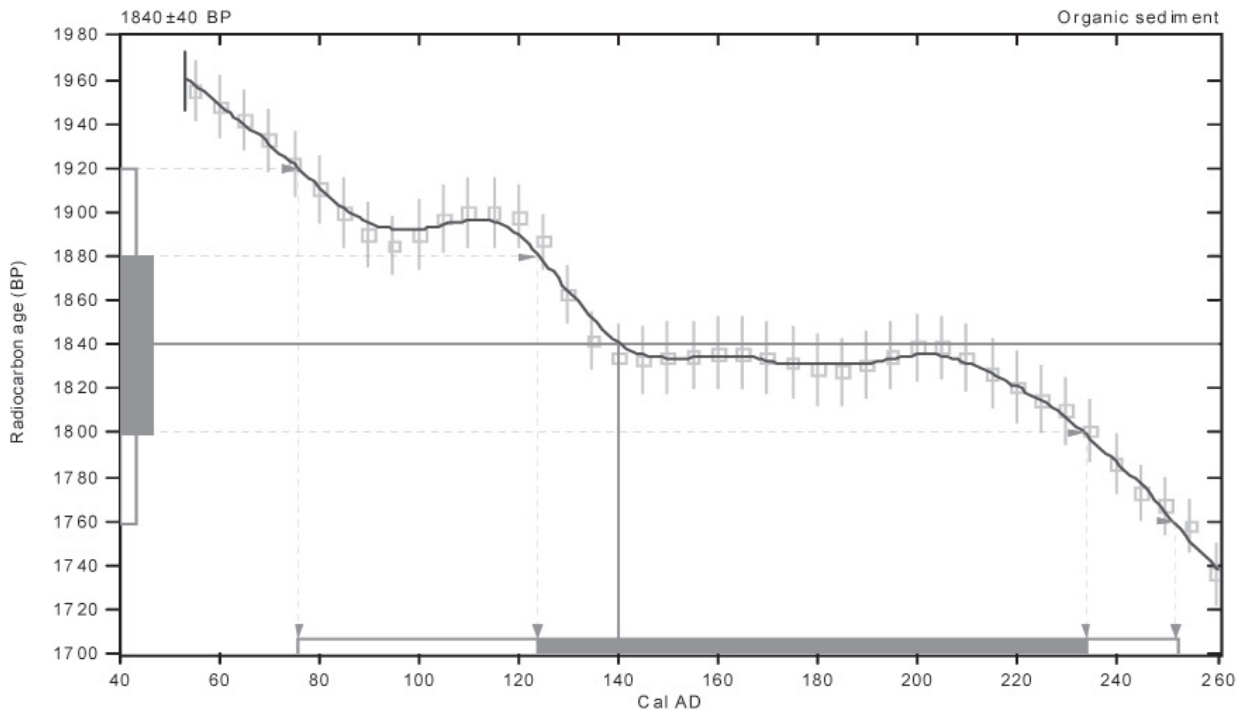
Conventional radiocarbon age: 1840±40 BP

**2 Sigma calibrated result: Cal AD 80 to 250 (Cal BP 1870 to 1700)
(95% probability)**

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 140 (Cal BP 1810)

**1 Sigma calibrated result: Cal AD 120 to 230 (Cal BP 1830 to 1720)
(68% probability)**



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-21.1:lab. mult=1)

Laboratory number: **Beta-235818**

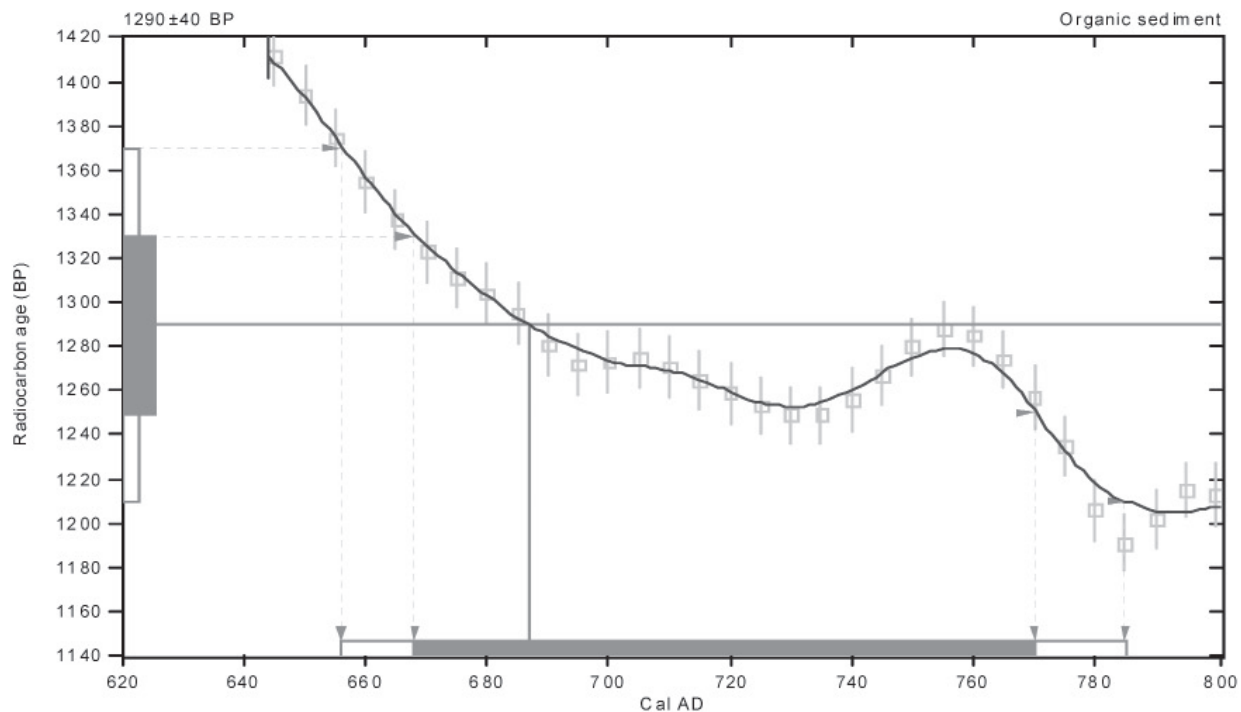
Conventional radiocarbon age: **1290±40 BP**

2 Sigma calibrated result: Cal AD 660 to 780 (Cal BP 1290 to 1160)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 690 (Cal BP 1260)

1 Sigma calibrated result: Cal AD 670 to 770 (Cal BP 1280 to 1180)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-22.1:lab. mult=1)

Laboratory number: **Beta-235820**

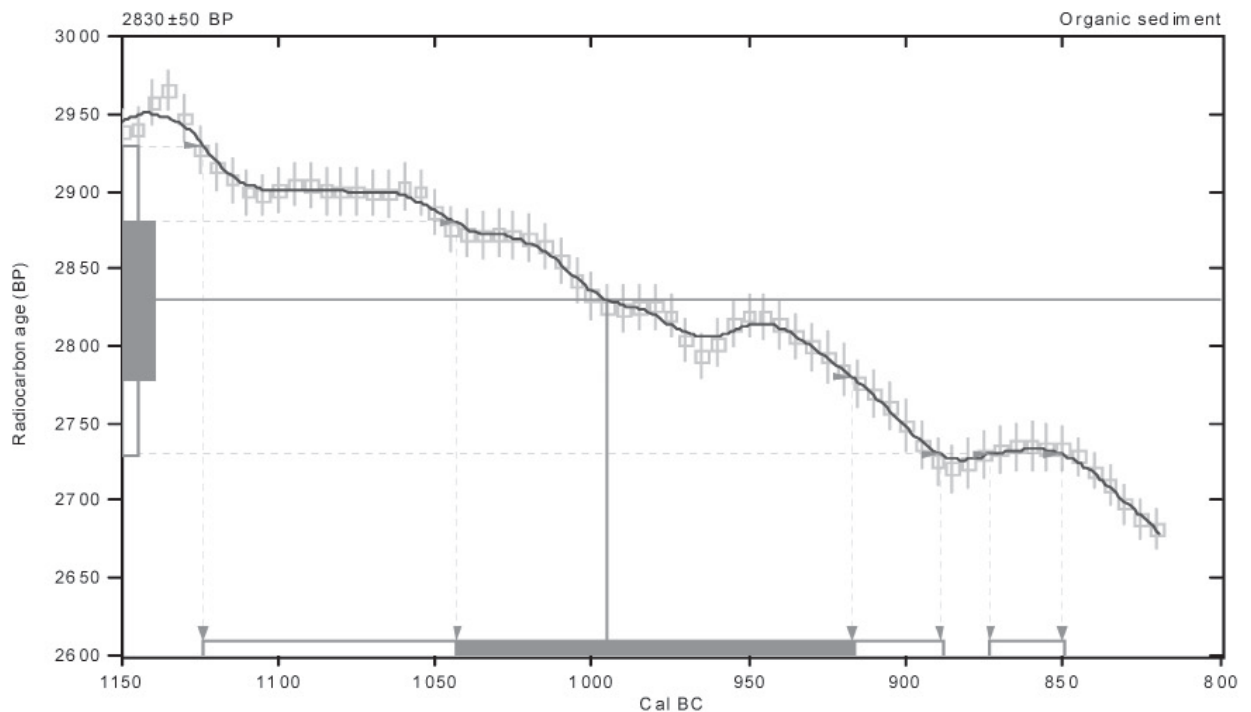
Conventional radiocarbon age: **2830±50 BP**

2 Sigma calibrated results: Cal BC 1120 to 890 (Cal BP 3070 to 2840) and
(95% probability) Cal BC 870 to 850 (Cal BP 2820 to 2800)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 1000 (Cal BP 2940)

1 Sigma calibrated result: Cal BC 1040 to 920 (Cal BP 2990 to 2870)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-25.8:lab. mult=1)

Laboratory number: **Beta-235821**

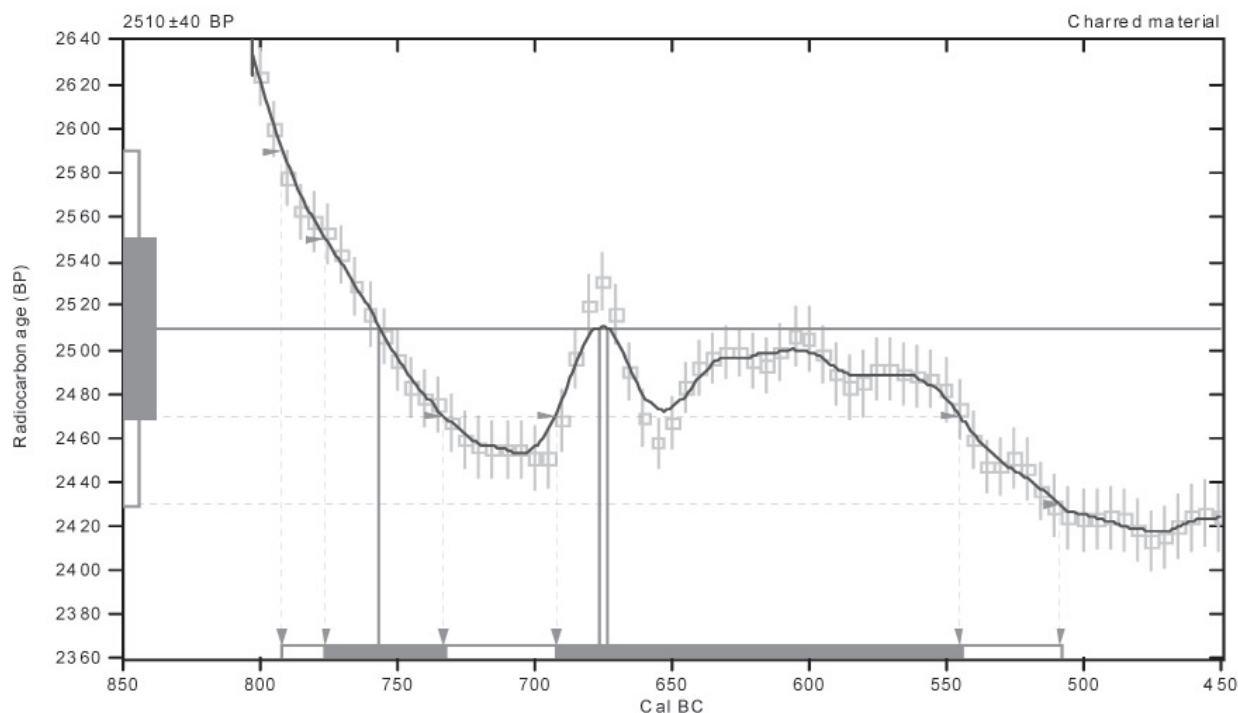
Conventional radiocarbon age: **2510±40 BP**

2 Sigma calibrated result: Cal BC 790 to 510 (Cal BP 2740 to 2460)
(95% probability)

Intercept data

Intercepts of radiocarbon age
with calibration curve: Cal BC 760 (Cal BP 2710) and
Cal BC 680 (Cal BP 2630) and
Cal BC 670 (Cal BP 2620)

1 Sigma calibrated results: Cal BC 780 to 730 (Cal BP 2730 to 2680) and
(68% probability) **Cal BC 690 to 540 (Cal BP 2640 to 2500)**



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-21.1:lab. mult=1)

Laboratory number: **Beta-235822**

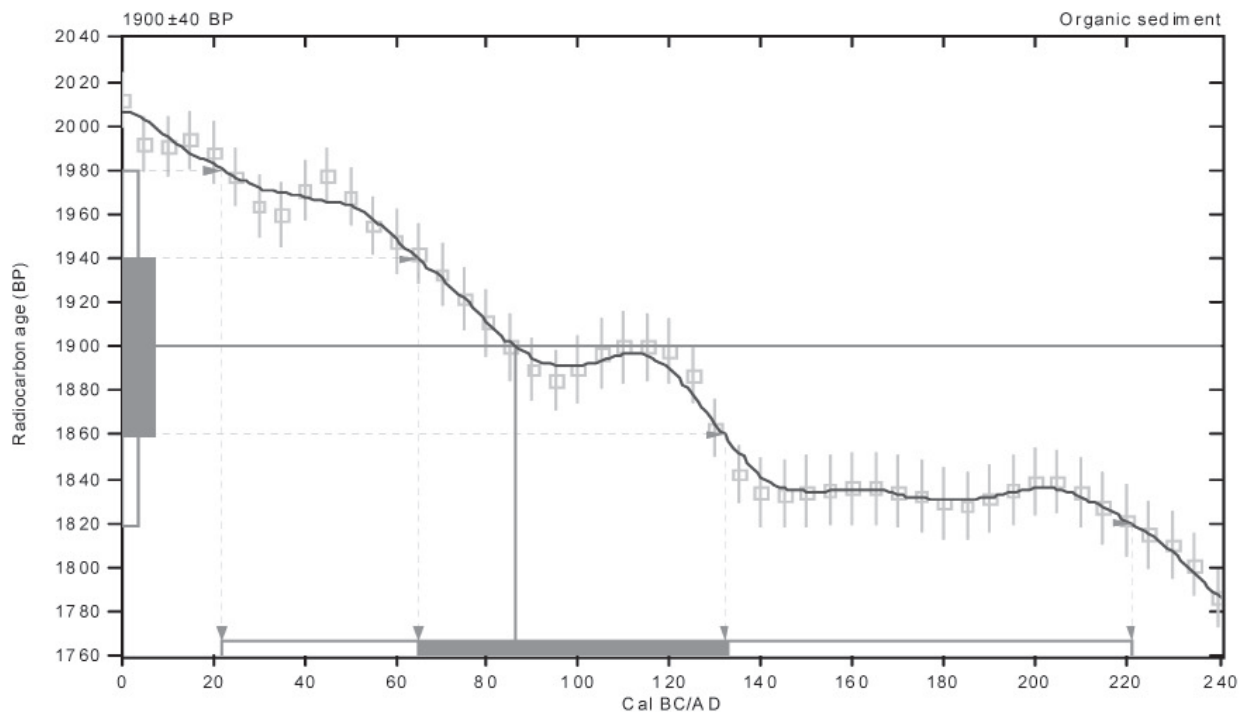
Conventional radiocarbon age: **1900±40 BP**

2 Sigma calibrated result: Cal AD 20 to 220 (Cal BP 1930 to 1730)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 90 (Cal BP 1860)**

1 Sigma calibrated result: Cal AD 60 to 130 (Cal BP 1880 to 1820)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.8:lab. mult=1)

Laboratory number: **Beta-235823**

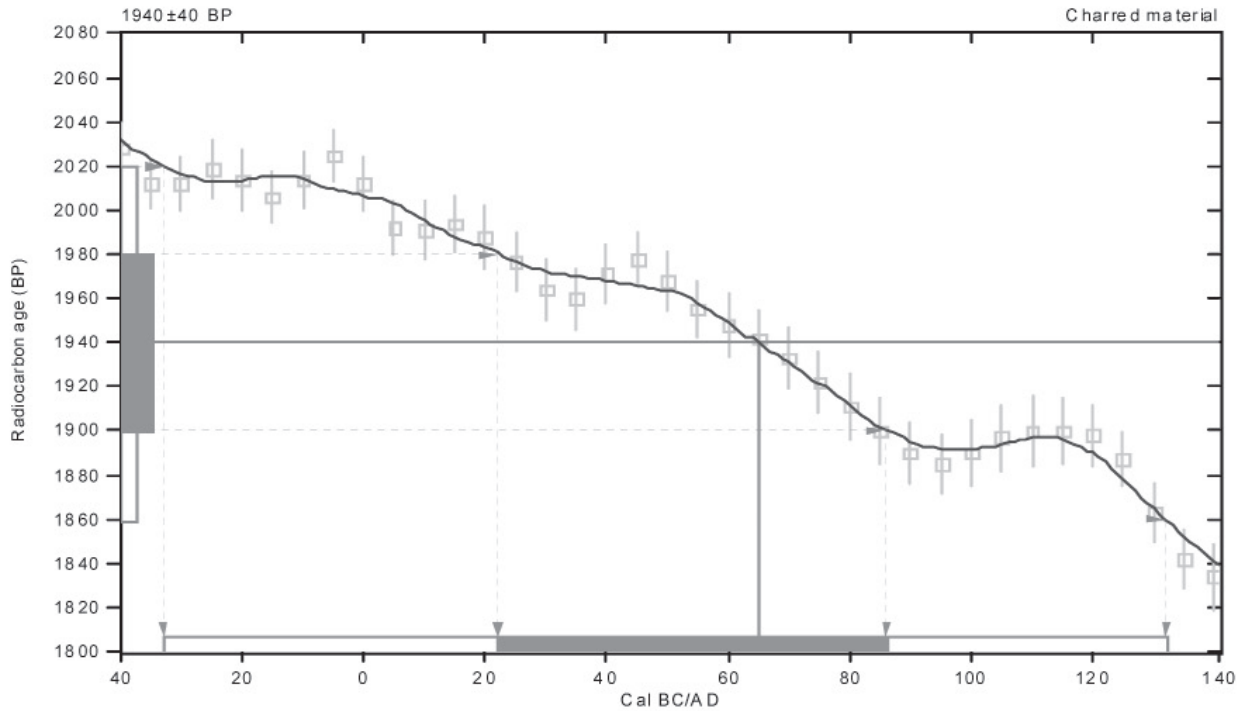
Conventional radiocarbon age: **1940±40 BP**

2 Sigma calibrated result: Cal BC 30 to Cal AD 130 (Cal BP 1980 to 1820)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 60 (Cal BP 1880)**

1 Sigma calibrated result: Cal AD 20 to 90 (Cal BP 1930 to 1860)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-18;lab.mult=1)

Laboratory number: **Beta-235824**

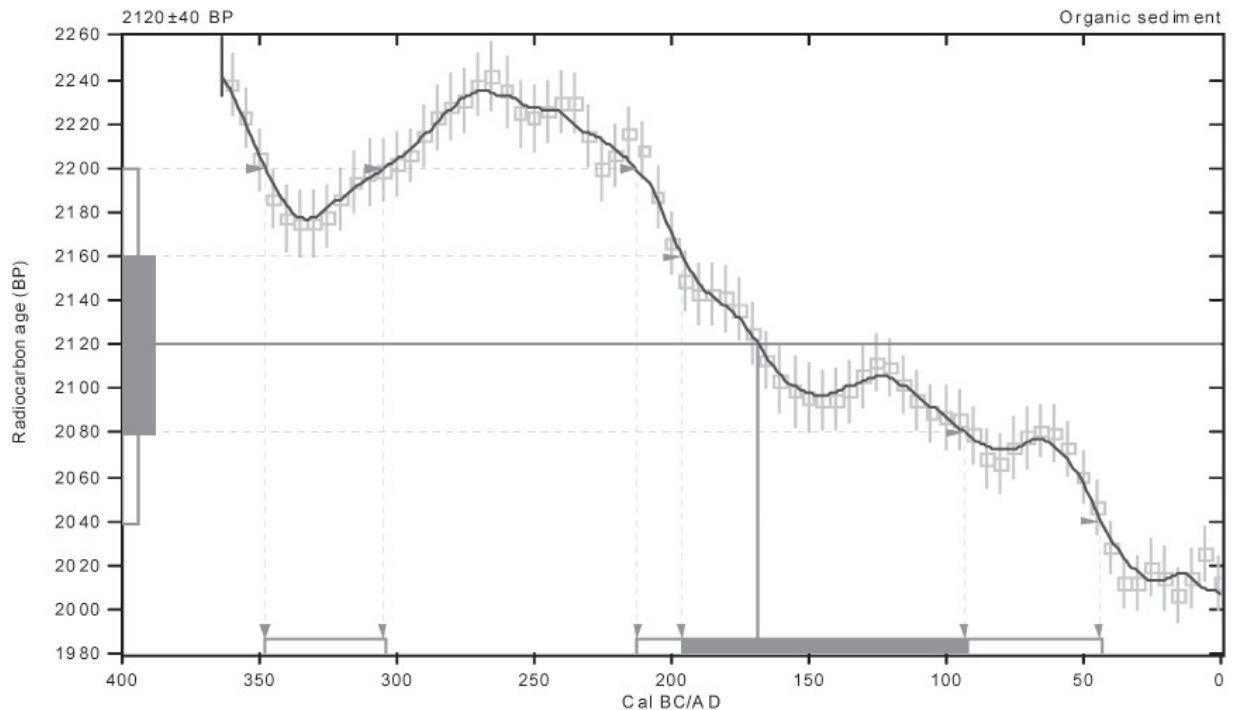
Conventional radiocarbon age: **2120±40 BP**

**2 Sigma calibrated results: Cal BC 350 to 300 (Cal BP 2300 to 2260) and
(95% probability) Cal BC 210 to 40 (Cal BP 2160 to 1990)**

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 170 (Cal BP 2120)**

**1 Sigma calibrated result: Cal BC 200 to 90 (Cal BP 2150 to 2040)
(68% probability)**



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-18.3:lab. mult=1)

Laboratory number: **Beta-235825**

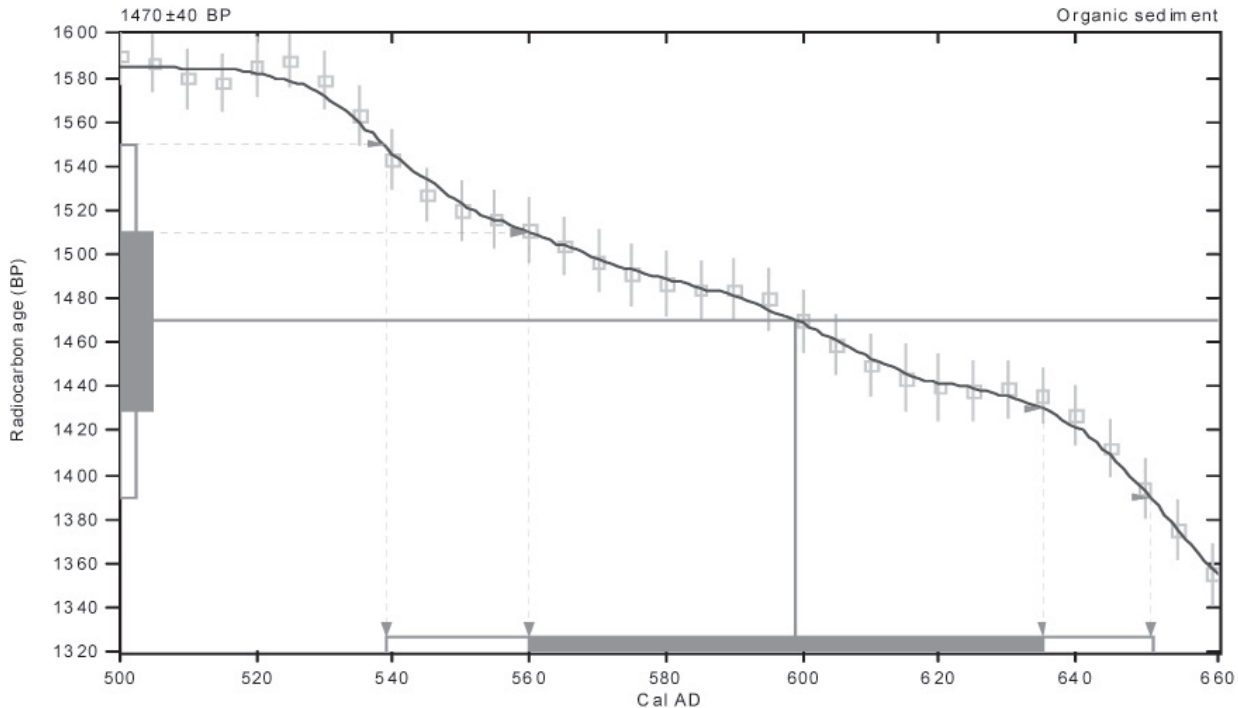
Conventional radiocarbon age: **1470±40 BP**

2 Sigma calibrated result: **Cal AD 540 to 650 (Cal BP 1410 to 1300)**
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 600 (Cal BP 1350)**

1 Sigma calibrated result: **Cal AD 560 to 640 (Cal BP 1390 to 1320)**
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.7:lab. mult=1)

Laboratory number: **Beta-235826**

Conventional radiocarbon age: **3300±40 BP**

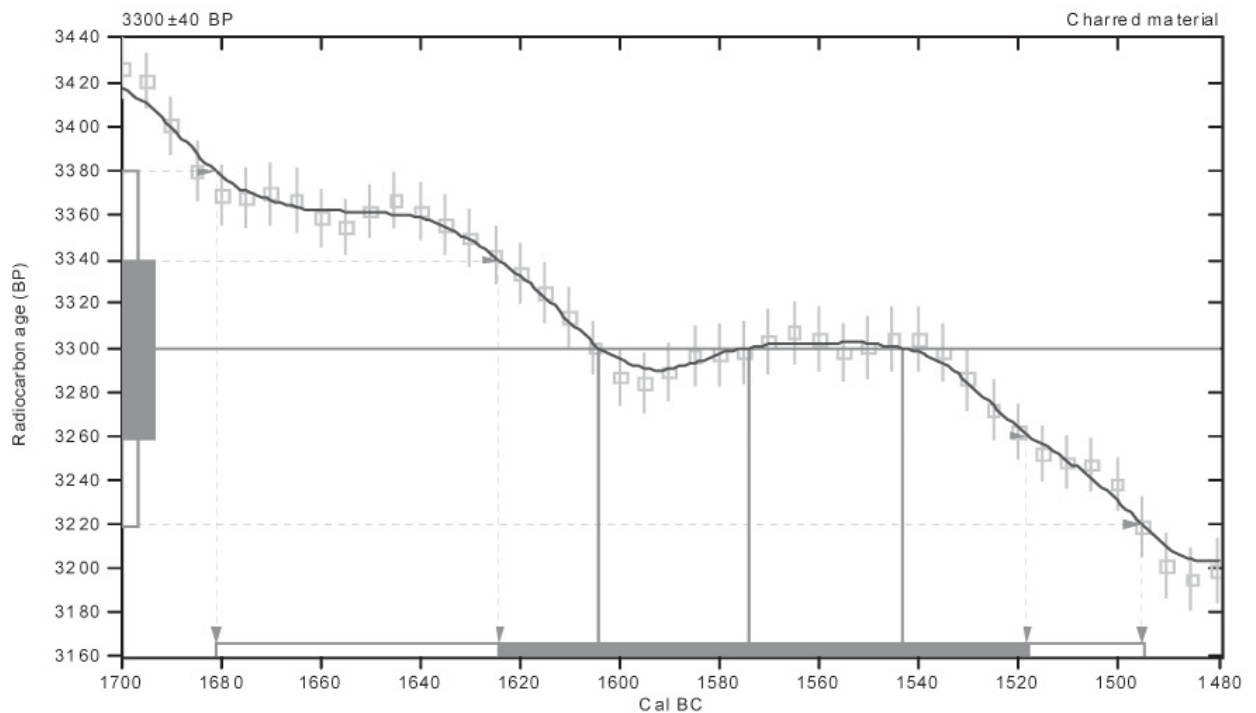
2 Sigma calibrated result: Cal BC 1680 to 1500 (Cal BP 3630 to 3440)
(95% probability)

Intercept data

Intercepts of radiocarbon age

with calibration curve: Cal BC 1600 (Cal BP 3550) and
Cal BC 1570 (Cal BP 3520) and
Cal BC 1540 (Cal BP 3490)

1 Sigma calibrated result: Cal BC 1620 to 1520 (Cal BP 3570 to 3470)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-19.1:lab. mult=1)

Laboratory number: **Beta-235827**

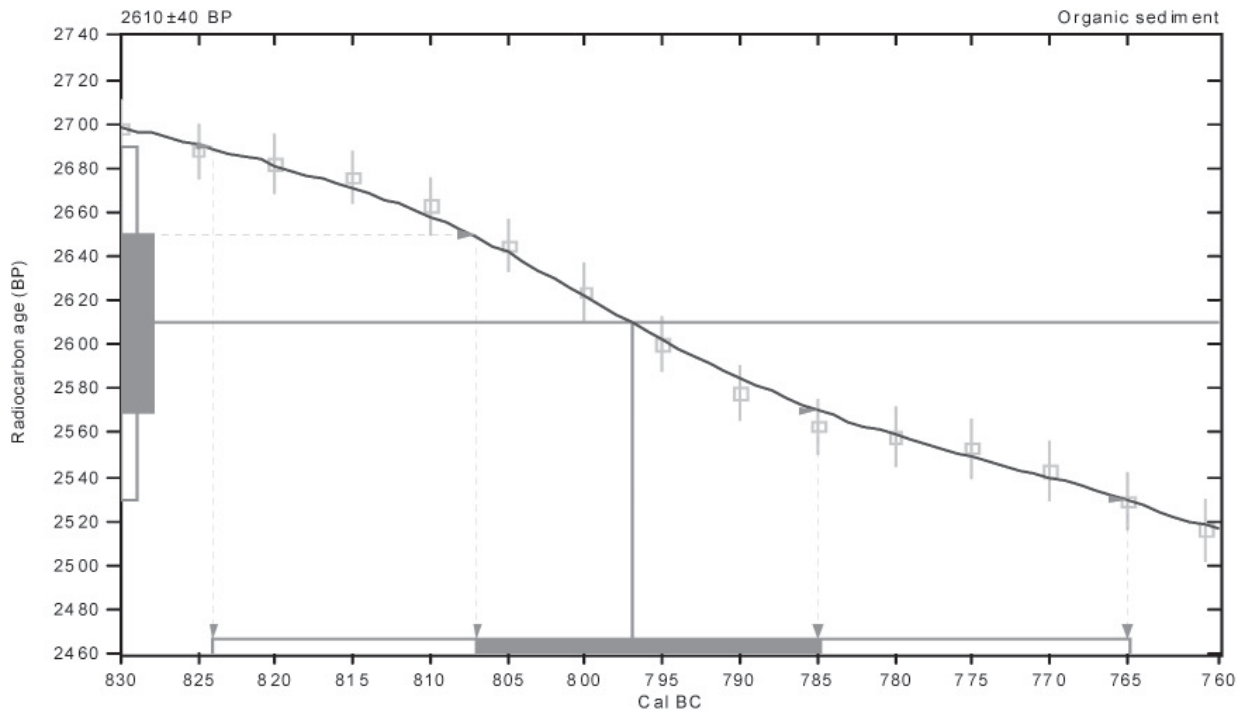
Conventional radiocarbon age: **2610±40 BP**

2 Sigma calibrated result: Cal BC 820 to 760 (Cal BP 2770 to 2720)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 800 (Cal BP 2750)

1 Sigma calibrated result: Cal BC 810 to 780 (Cal BP 2760 to 2740)
(68% probability)



References:

- Database used*
INTCAL04
- Calibration Database*
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.3:lab. mult=1)

Laboratory number: **Beta-235828**

Conventional radiocarbon age: **2470±40 BP**

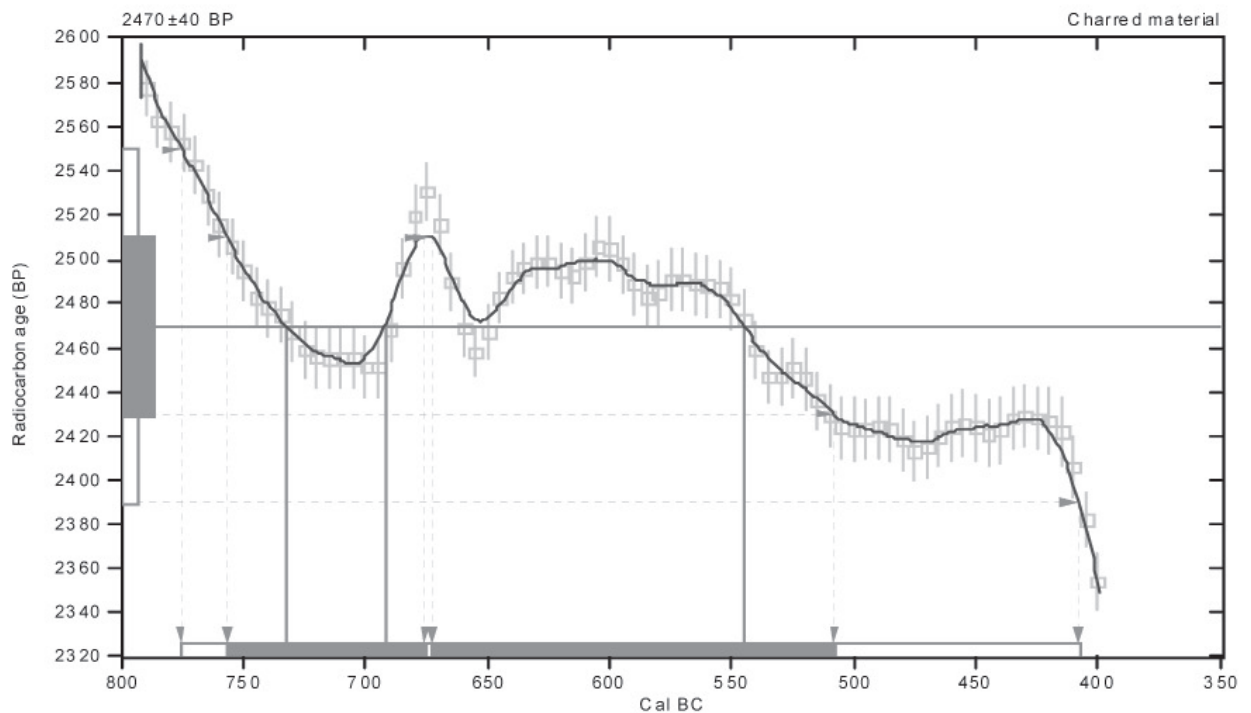
2 Sigma calibrated result: **Cal BC 780 to 410 (Cal BP 2730 to 2360)**
(95% probability)

Intercept data

Intercepts of radiocarbon age
with calibration curve:

Cal BC 730 (Cal BP 2680) and
Cal BC 690 (Cal BP 2640) and
Cal BC 540 (Cal BP 2500)

1 Sigma calibrated results: **Cal BC 760 to 680 (Cal BP 2710 to 2630) and**
(68% probability) **Cal BC 670 to 510 (Cal BP 2620 to 2460)**



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-25.7;lab. mult=1)

Laboratory number: **Beta-235829**

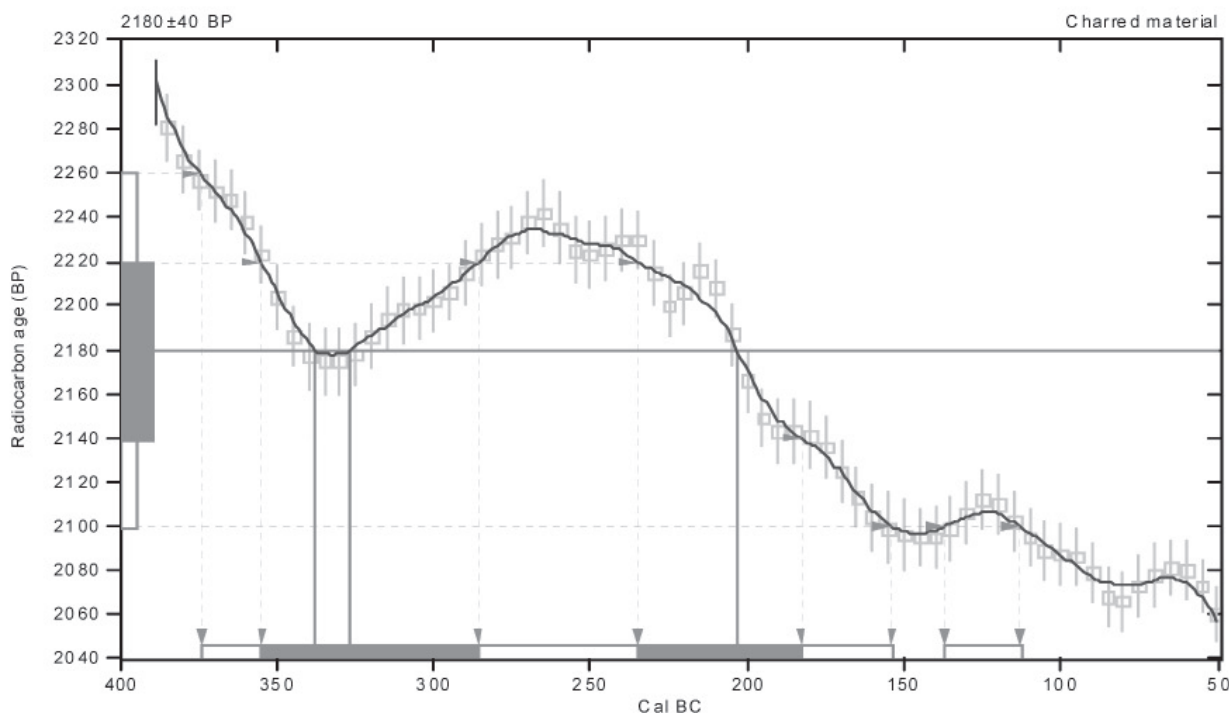
Conventional radiocarbon age: **2180±40 BP**

2 Sigma calibrated results: Cal BC 370 to 150 (Cal BP 2320 to 2100) and
(95% probability) Cal BC 140 to 110 (Cal BP 2090 to 2060)

Intercept data

Intercepts of radiocarbon age
with calibration curve: Cal BC 340 (Cal BP 2290) and
Cal BC 330 (Cal BP 2280) and
Cal BC 200 (Cal BP 2150)

1 Sigma calibrated results: Cal BC 360 to 290 (Cal BP 2300 to 2240) and
(68% probability) Cal BC 240 to 180 (Cal BP 2180 to 2130)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.9:lab. mult=1)

Laboratory number: **Beta-235830**

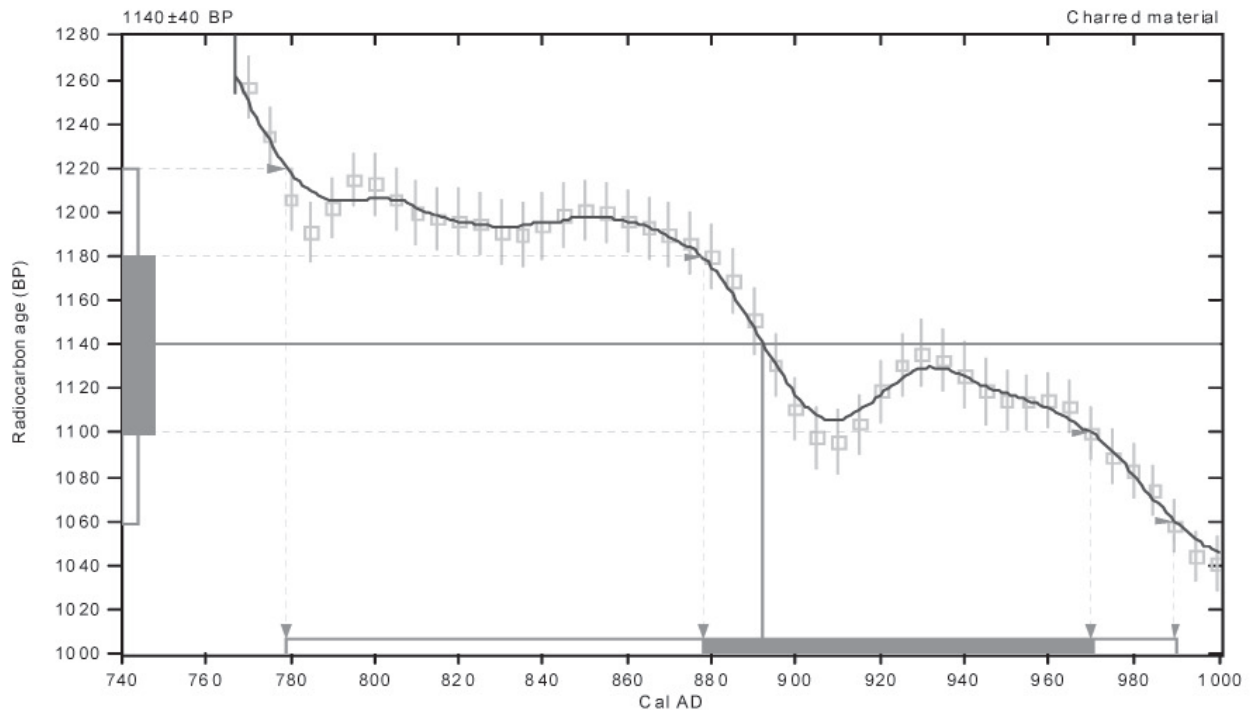
Conventional radiocarbon age: **1140±40 BP**

2 Sigma calibrated result: Cal AD 780 to 990 (Cal BP 1170 to 960)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 890 (Cal BP 1060)

1 Sigma calibrated result: Cal AD 880 to 970 (Cal BP 1070 to 980)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-18.8:lab. mult=1)

Laboratory number: **Beta-235831**

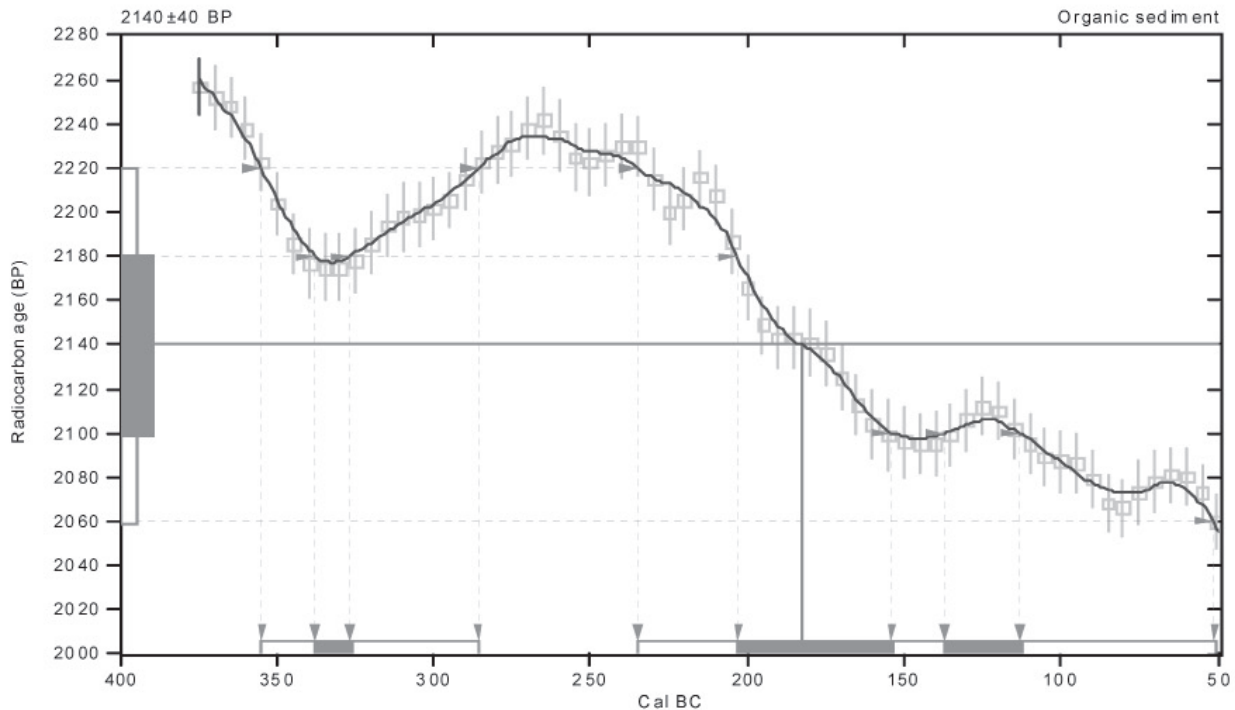
Conventional radiocarbon age: **2140±40 BP**

2 Sigma calibrated results: Cal BC 360 to 290 (Cal BP 2300 to 2240) and
(95% probability) Cal BC 240 to 50 (Cal BP 2180 to 2000)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 180 (Cal BP 2130)

1 Sigma calibrated results: Cal BC 340 to 330 (Cal BP 2290 to 2280) and
(68% probability) Cal BC 200 to 150 (Cal BP 2150 to 2100) and
Cal BC 140 to 110 (Cal BP 2090 to 2060)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-19.3:lab. mult=1)

Laboratory number: **Beta-235832**

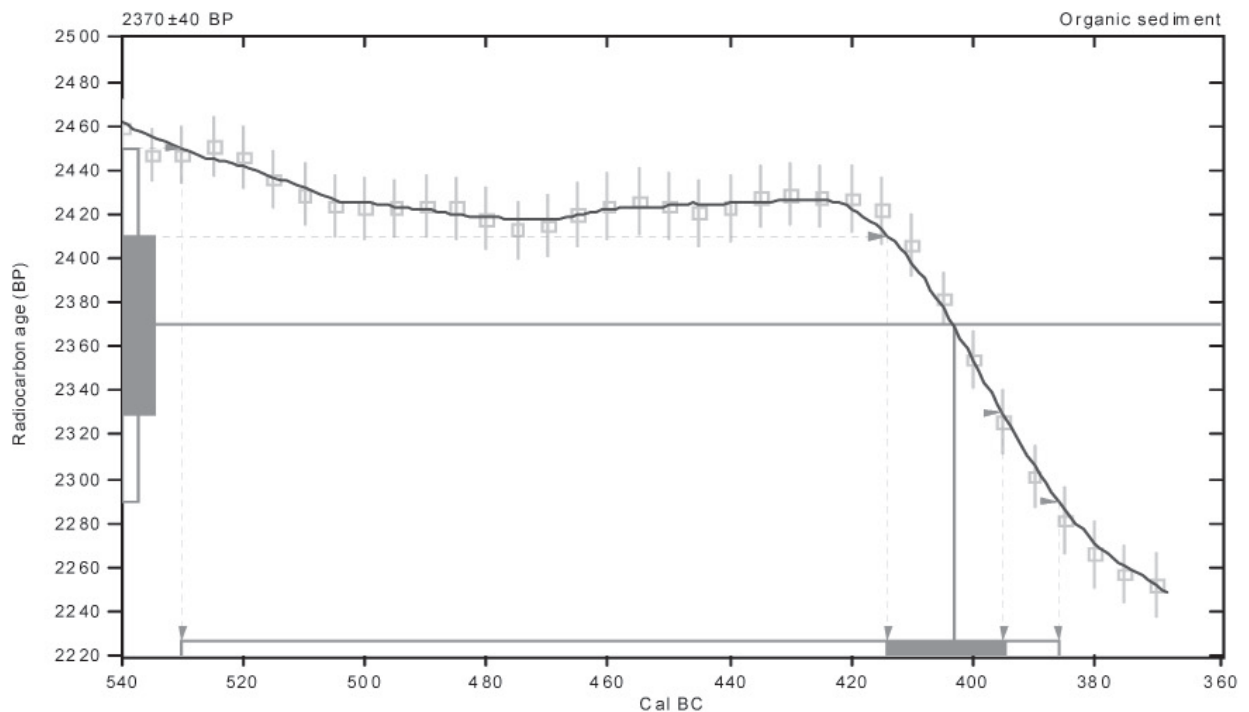
Conventional radiocarbon age: **2370±40 BP**

2 Sigma calibrated result: Cal BC 530 to 390 (Cal BP 2480 to 2340)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 400 (Cal BP 2350)**

1 Sigma calibrated result: Cal BC 410 to 400 (Cal BP 2360 to 2340)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-22.6:lab. mult=1)

Laboratory number: **Beta-235833**

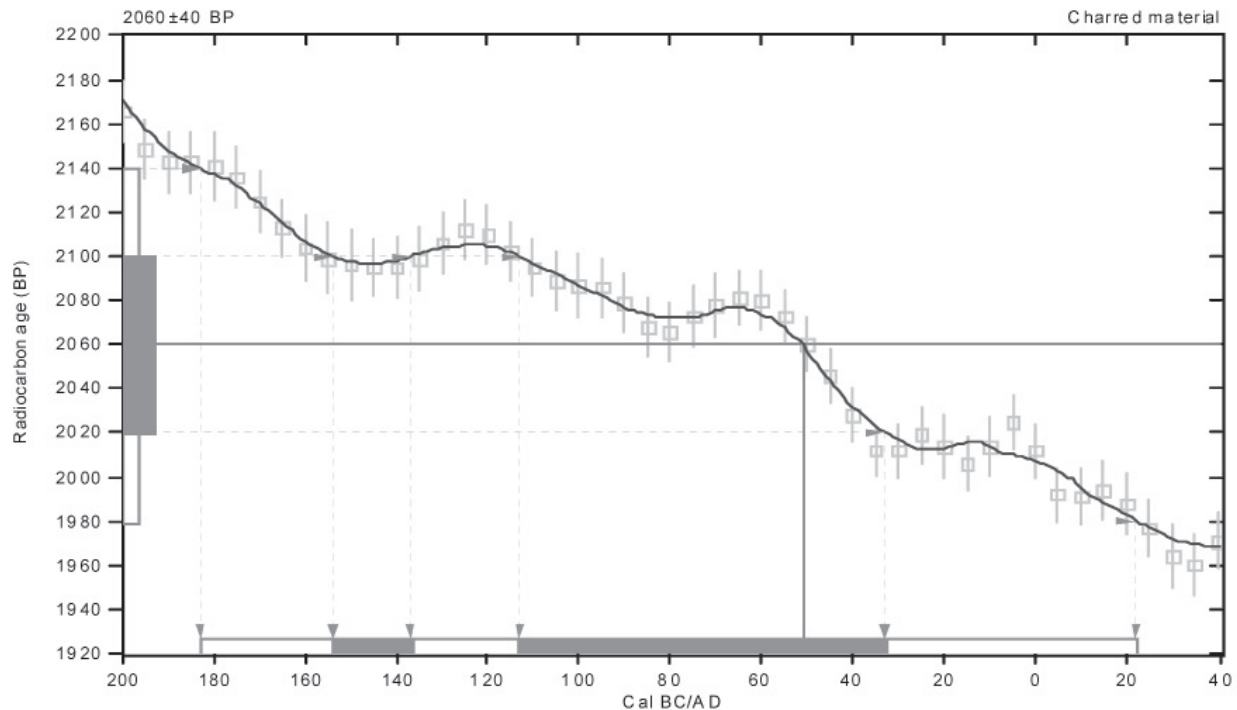
Conventional radiocarbon age: **2060±40 BP**

2 Sigma calibrated result: Cal BC 180 to Cal AD 20 (Cal BP 2130 to 1930)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 50 (Cal BP 2000)**

1 Sigma calibrated results: Cal BC 150 to 140 (Cal BP 2100 to 2090) and
Cal BC 110 to 30 (Cal BP 2060 to 1980)
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-24.1:lab. mult=1)

Laboratory number: **Beta-235834**

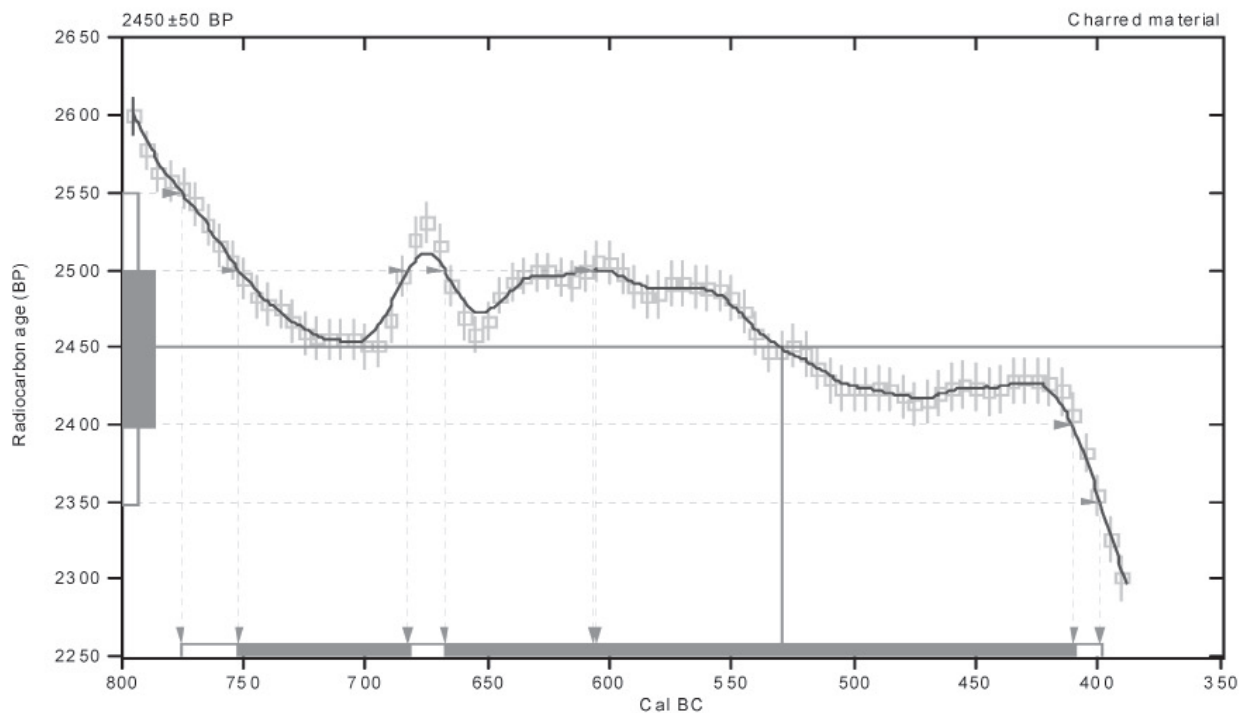
Conventional radiocarbon age: **2450±50 BP**

2 Sigma calibrated result: Cal BC 780 to 400 (Cal BP 2730 to 2350)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 530 (Cal BP 2480)

1 Sigma calibrated results: Cal BC 750 to 680 (Cal BP 2700 to 2630) and
(68% probability) Cal BC 670 to 610 (Cal BP 2620 to 2560) and
Cal BC 600 to 410 (Cal BP 2560 to 2360)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.9;lab. mult=1)

Laboratory number: **Beta-235835**

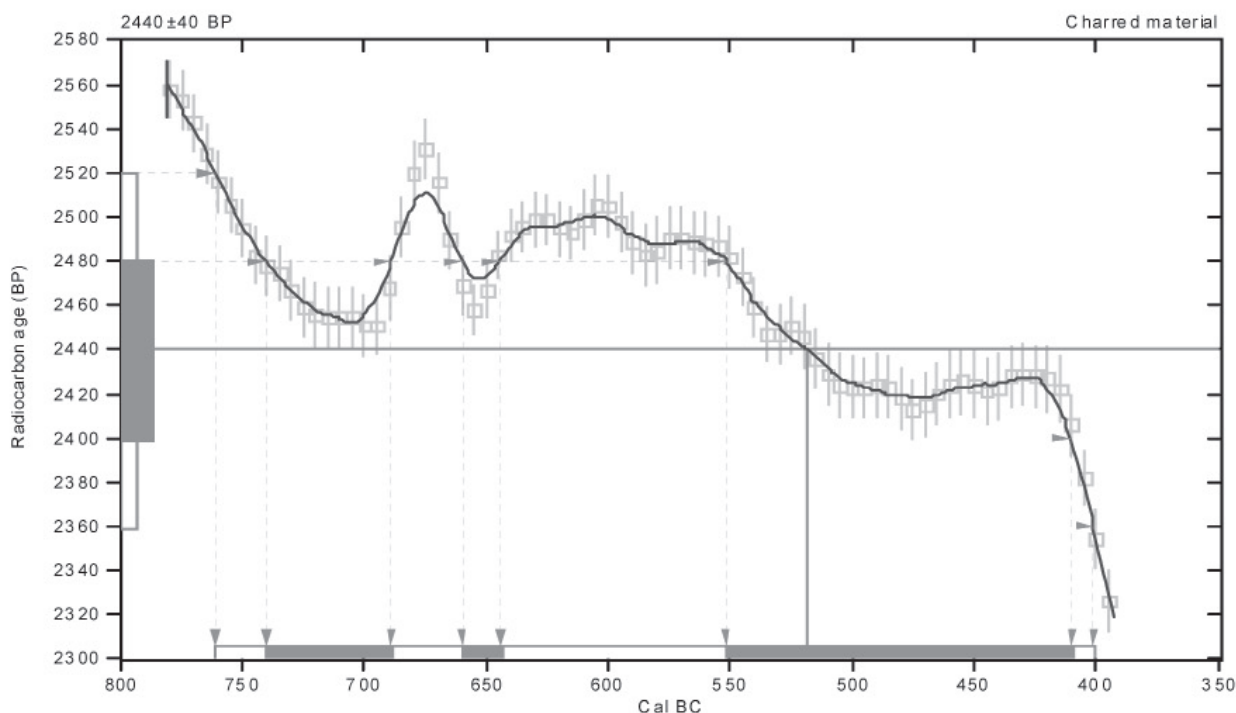
Conventional radiocarbon age: **2440±40 BP**

2 Sigma calibrated result: Cal BC 760 to 400 (Cal BP 2710 to 2350)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 520 (Cal BP 2470)

1 Sigma calibrated results: Cal BC 740 to 690 (Cal BP 2690 to 2640) and
(68% probability) Cal BC 660 to 640 (Cal BP 2610 to 2590) and
Cal BC 550 to 410 (Cal BP 2500 to 2360)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-23.4:lab. mult=1)

Laboratory number: **Beta-235836**

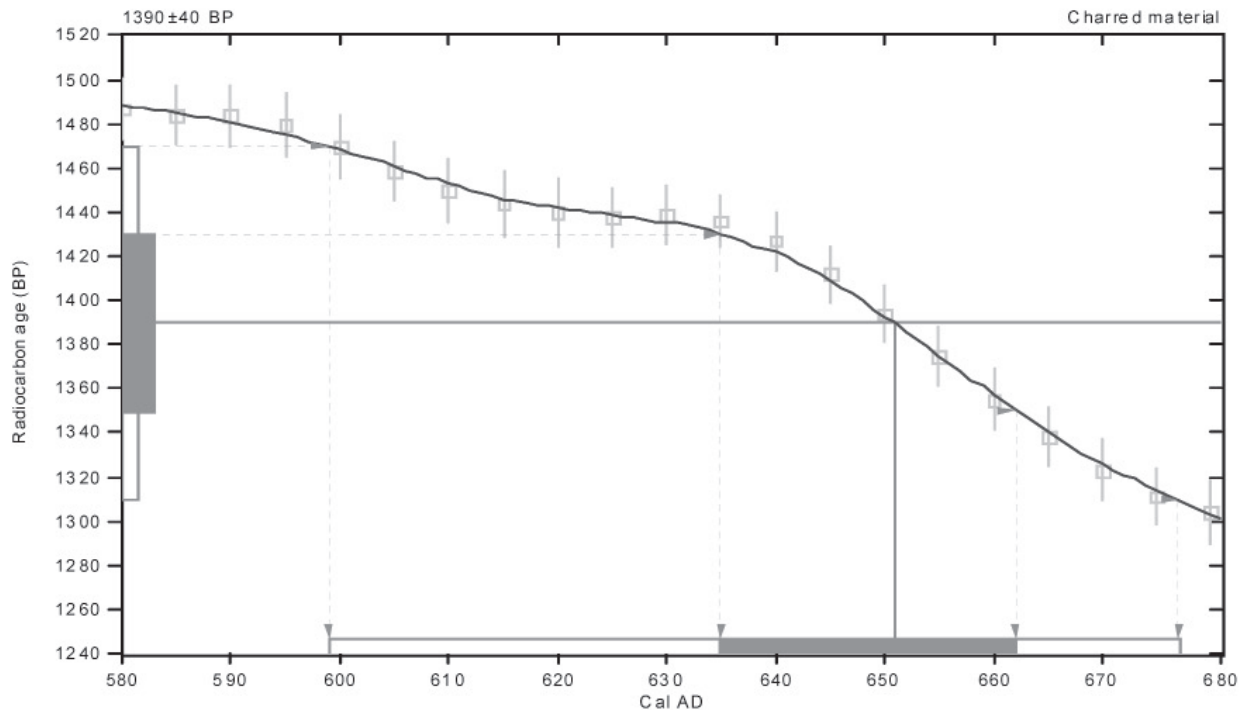
Conventional radiocarbon age: **1390±40 BP**

2 Sigma calibrated result: Cal AD 600 to 680 (Cal BP 1350 to 1270)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 650 (Cal BP 1300)**

1 Sigma calibrated result: Cal AD 640 to 660 (Cal BP 1320 to 1290)
(68% probability)



References:

Database used

INTCAL04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-25:lab. mult=1)

Laboratory number: **Beta-235837**

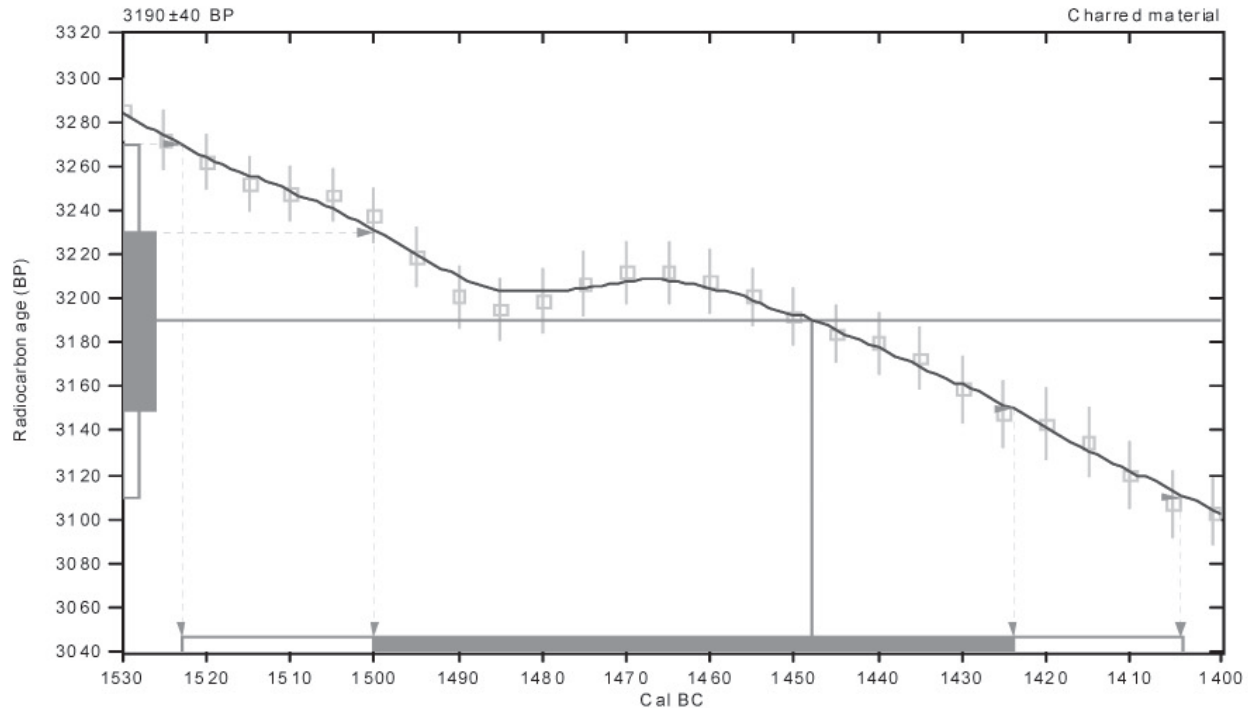
Conventional radiocarbon age: **3190±40 BP**

2 Sigma calibrated result: **Cal BC 1520 to 1400 (Cal BP 3470 to 3350)**
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 1450 (Cal BP 3400)**

1 Sigma calibrated result: **Cal BC 1500 to 1420 (Cal BP 3450 to 3370)**
(68% probability)



References:

- Database used*
INTCAL04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.4;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: Beta-235838

Conventional radiocarbon age: 1390±40 BP

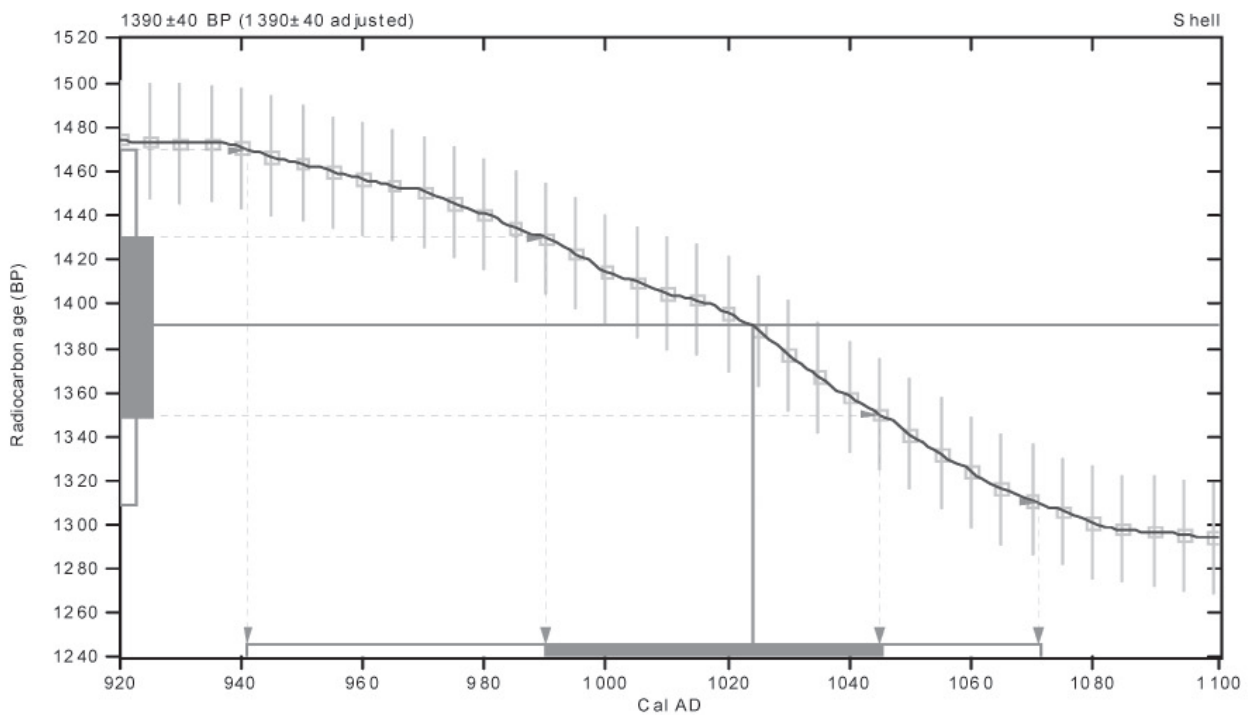
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 940 to 1070 (Cal BP 1010 to 880)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 1020 (Cal BP 930)

1 Sigma calibrated result: Cal AD 990 to 1040 (Cal BP 960 to 900)
(68% probability)



References:

Database used
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.4:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: Beta-235839

Conventional radiocarbon age: 2090±40 BP

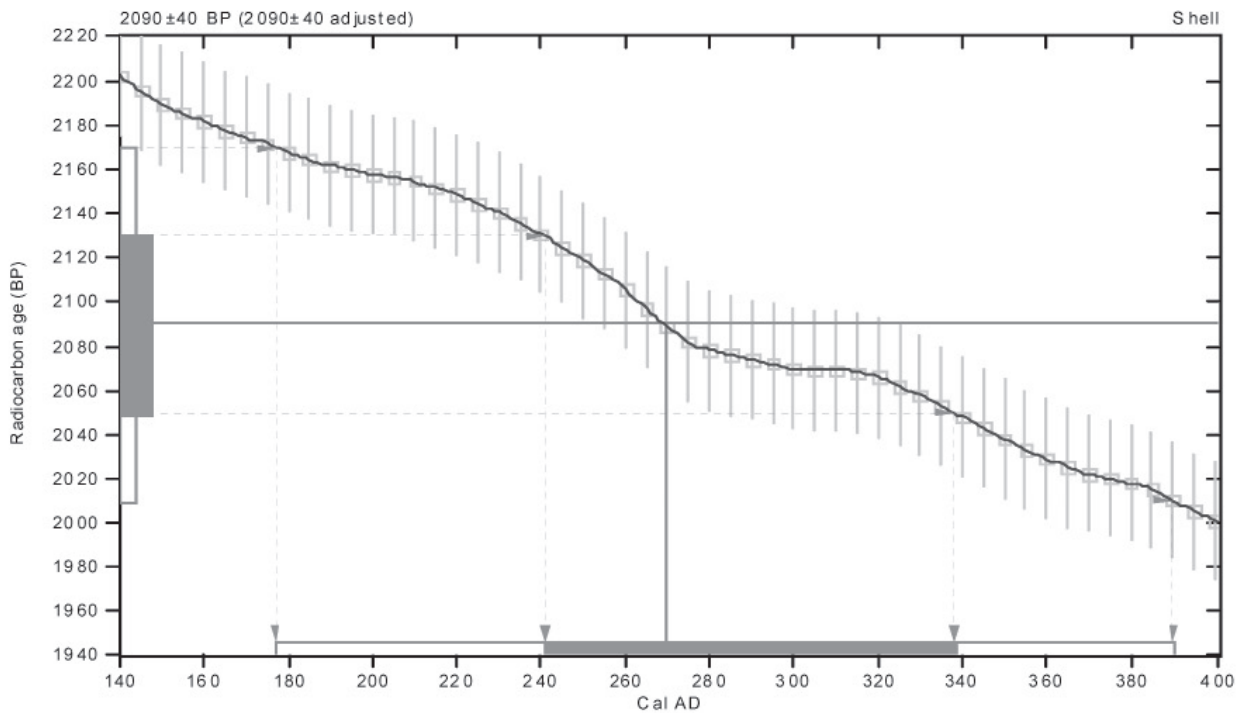
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 180 to 390 (Cal BP 1770 to 1560)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 270 (Cal BP 1680)

1 Sigma calibrated result: Cal AD 240 to 340 (Cal BP 1710 to 1610)
(68% probability)



References:

Database used

MARINE 04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.8:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: Beta-235841

Conventional radiocarbon age: 2100±40 BP

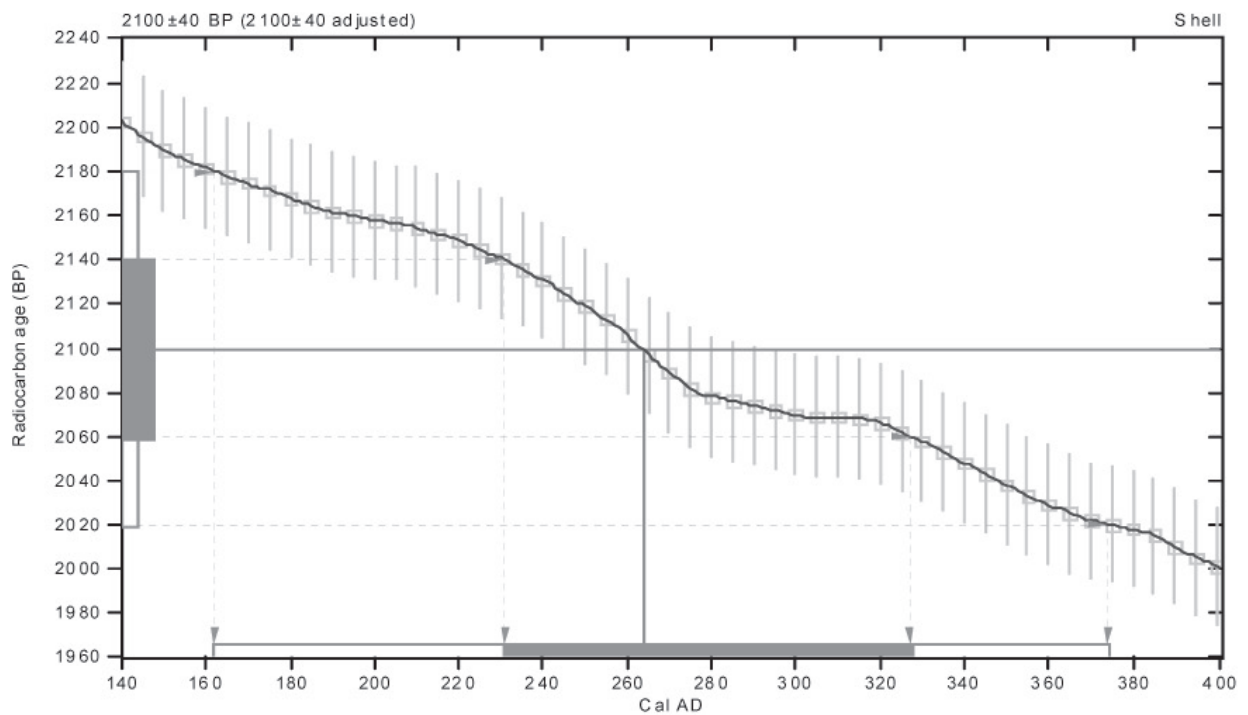
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 160 to 370 (Cal BP 1790 to 1580)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 260 (Cal BP 1690)

1 Sigma calibrated result: Cal AD 230 to 330 (Cal BP 1720 to 1620)
(68% probability)



References:

Database used

MARINE04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.2:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235842**

Conventional radiocarbon age: **2540±40 BP**

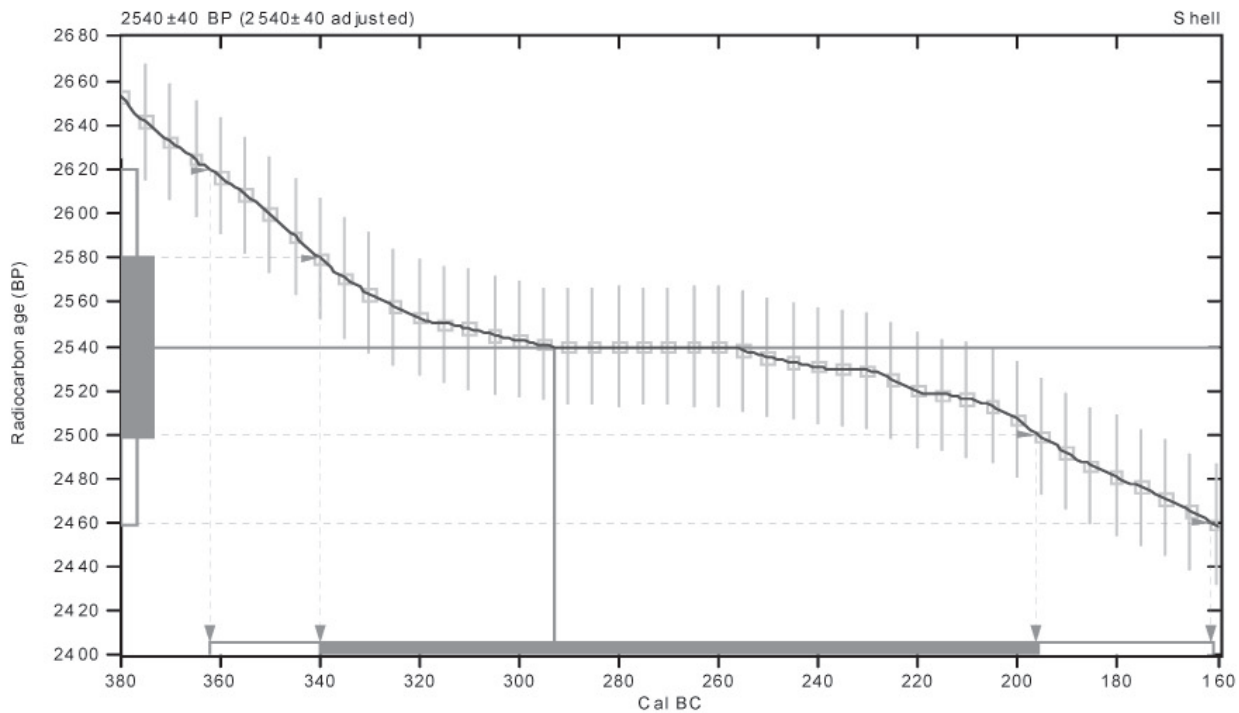
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 360 to 160 (Cal BP 2310 to 2110)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 290 (Cal BP 2240)

1 Sigma calibrated result: Cal BC 340 to 200 (Cal BP 2290 to 2150)
(68% probability)



References:

- Database used*
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
- Mathematics*
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.5:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: Beta-235843

Conventional radiocarbon age: 2200±40 BP

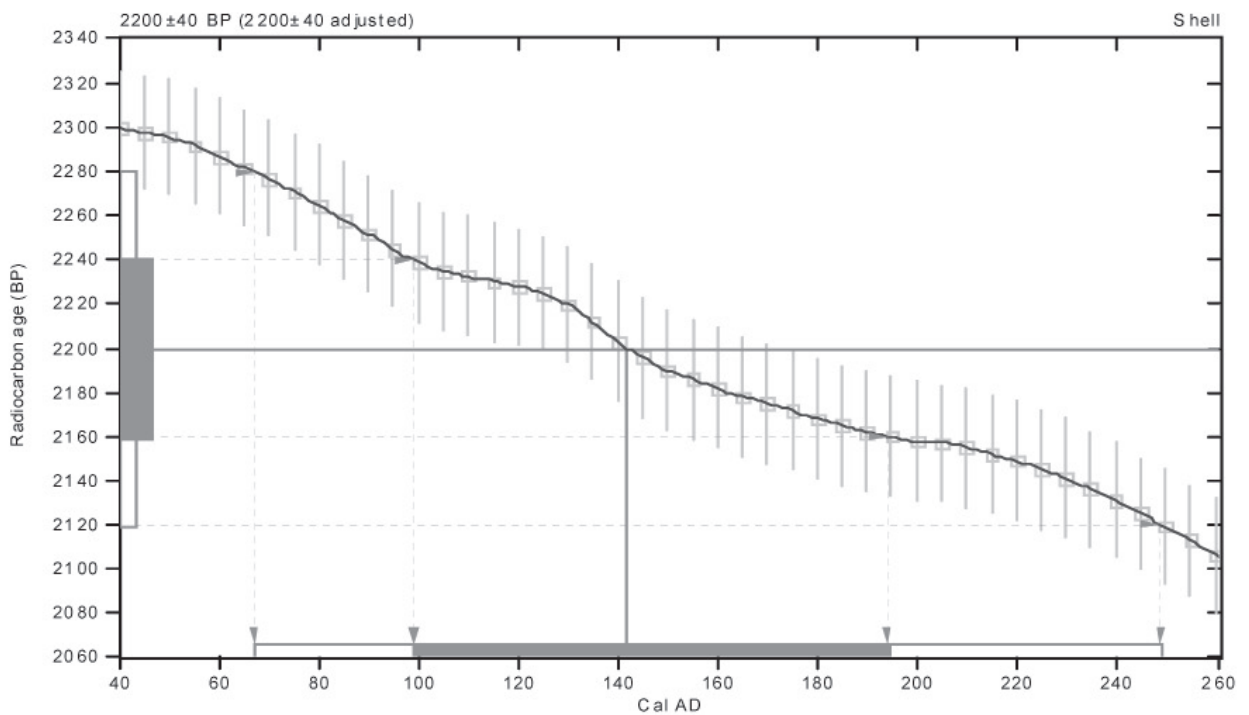
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 70 to 250 (Cal BP 1880 to 1700)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 140 (Cal BP 1810)

1 Sigma calibrated result: Cal AD 100 to 190 (Cal BP 1850 to 1760)
(68% probability)



References:

Database used
 MARINE04
Calibration Database
 INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.6:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235844**

Conventional radiocarbon age: **2850±40 BP**

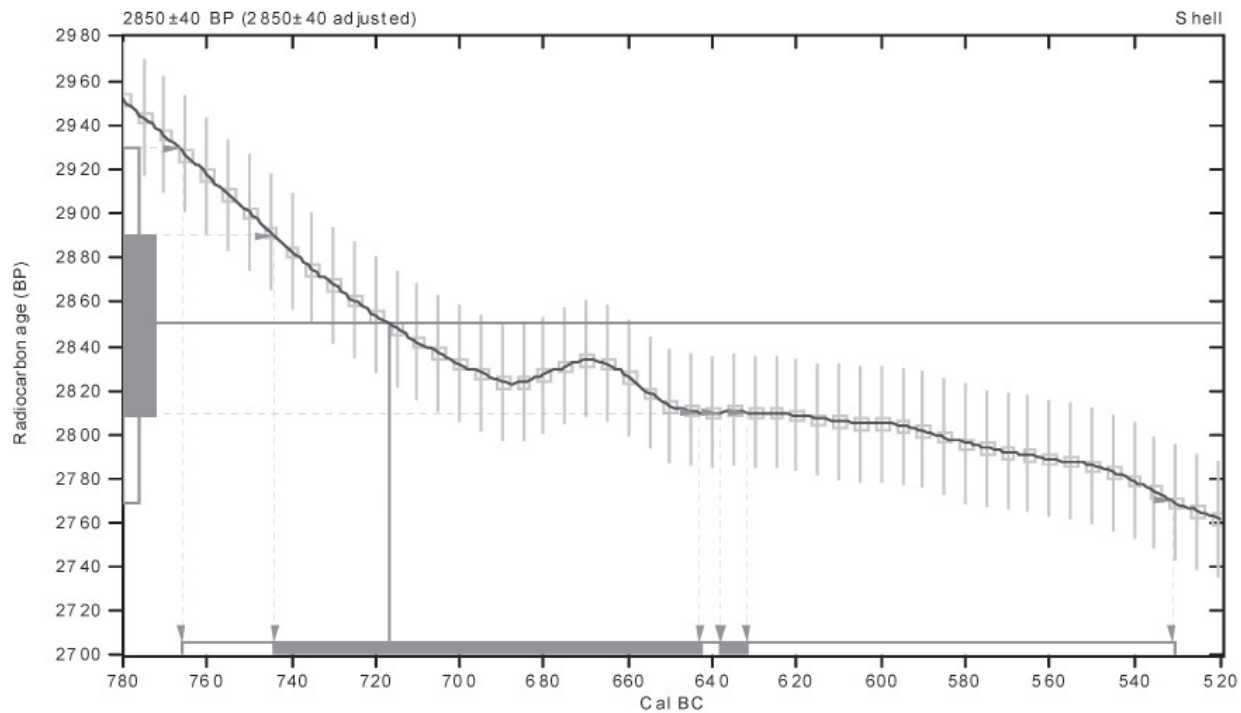
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 770 to 530 (Cal BP 2720 to 2480)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 720 (Cal BP 2670)

1 Sigma calibrated results: Cal BC 740 to 640 (Cal BP 2690 to 2590) and
(68% probability) Cal BC 640 to 630 (Cal BP 2590 to 2580)



References:

- Database used*
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.3;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: Beta-235845

Conventional radiocarbon age: 5540±40 BP

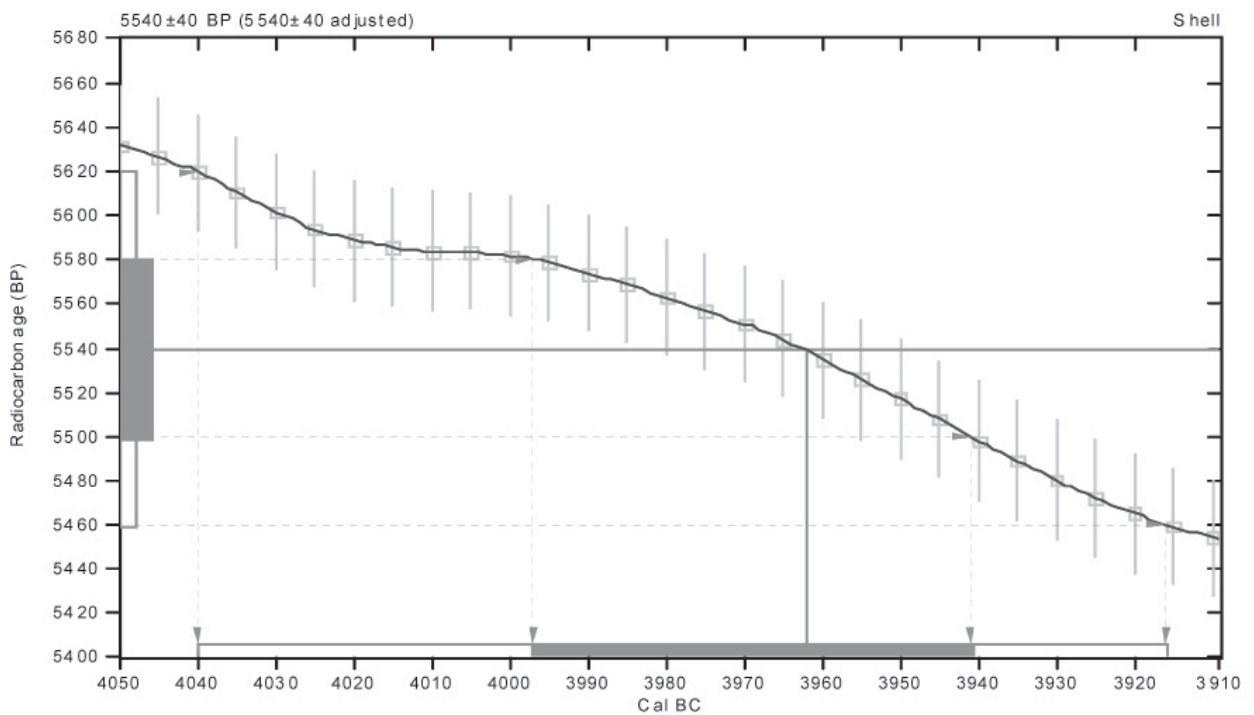
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 4040 to 3920 (Cal BP 5990 to 5870)
 (95% probability)

Intercept data

Intercept of radiocarbon age
 with calibration curve: Cal BC 3960 (Cal BP 5910)

1 Sigma calibrated result: Cal BC 4000 to 3940 (Cal BP 5950 to 5890)
 (68% probability)



References:

Database used
 MARINE04
Calibration Database
 INTCAL04 Radiocarbon Age Calibration
 IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
 A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.2;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: Beta-235846

Conventional radiocarbon age: 2110±40 BP

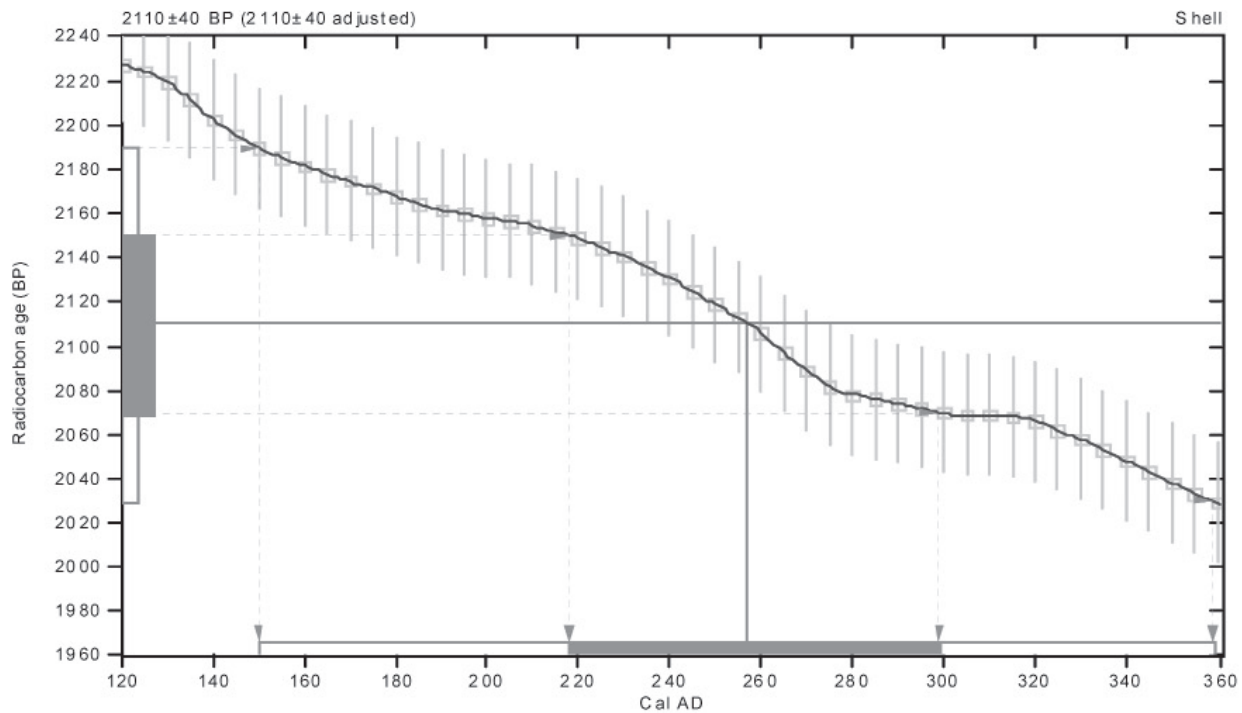
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 150 to 360 (Cal BP 1800 to 1590)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 260 (Cal BP 1690)

1 Sigma calibrated result: Cal AD 220 to 300 (Cal BP 1730 to 1650)
(68% probability)



References:

Data base used

MARINE 04

Calibration Data base

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=1:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235847**

Conventional radiocarbon age: **2590±40 BP**

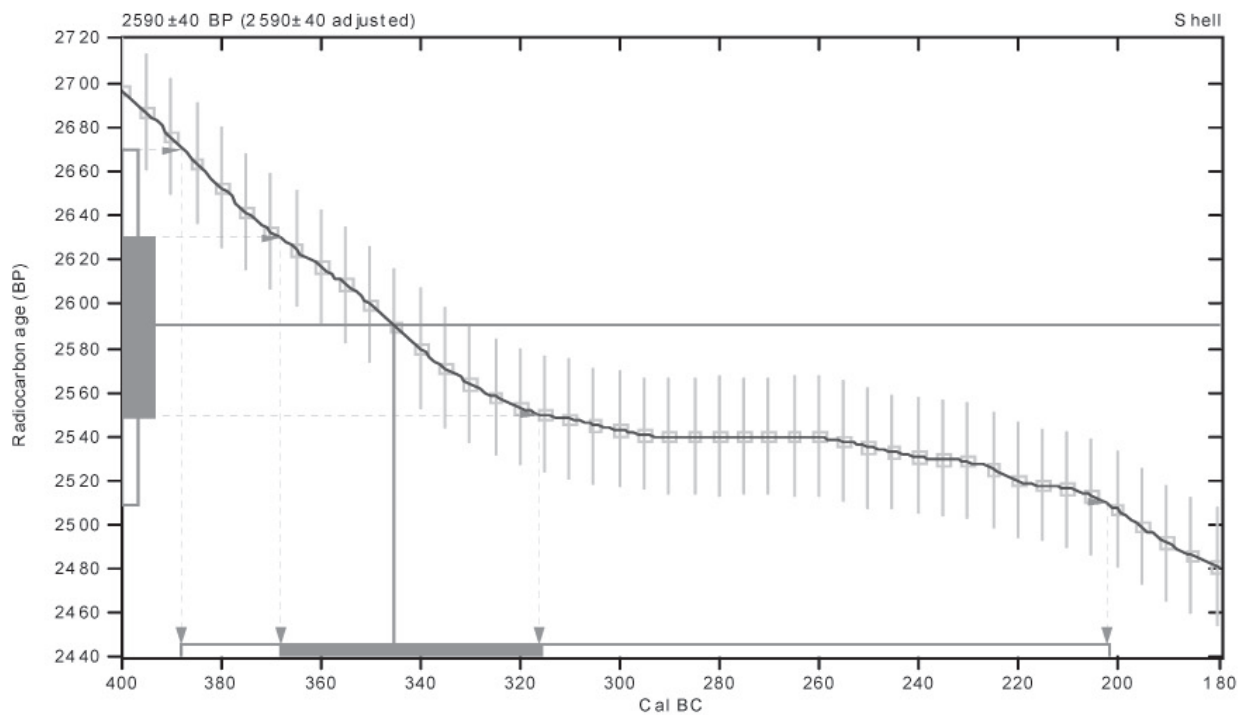
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 390 to 200 (Cal BP 2340 to 2150)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 340 (Cal BP 2300)

1 Sigma calibrated result: Cal BC 370 to 320 (Cal BP 2320 to 2270)
(68% probability)



References:

Database used

MARINE 04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.4:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235848**

Conventional radiocarbon age: **3090±40 BP**

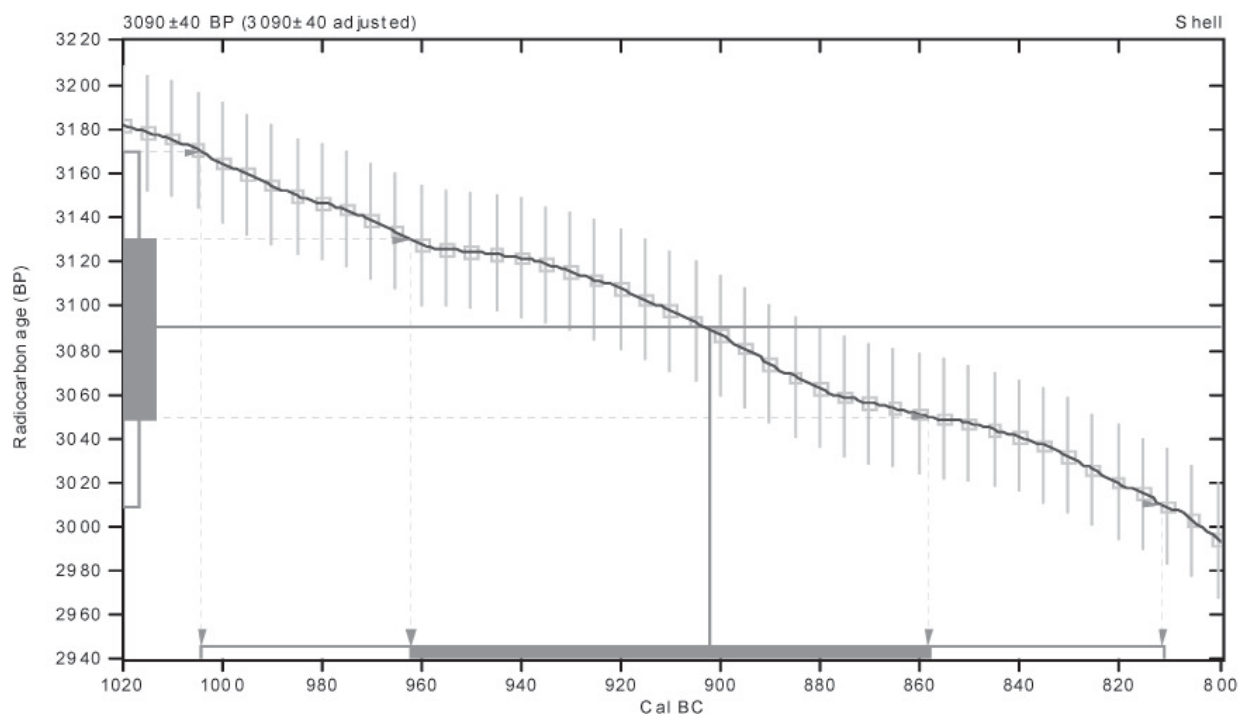
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 1000 to 810 (Cal BP 2950 to 2760)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 900 (Cal BP 2850)

1 Sigma calibrated result: Cal BC 960 to 860 (Cal BP 2910 to 2810)
(68% probability)



References:

Database used

MARINE04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.1;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: **Beta-235849**

Conventional radiocarbon age: **2030±40 BP**

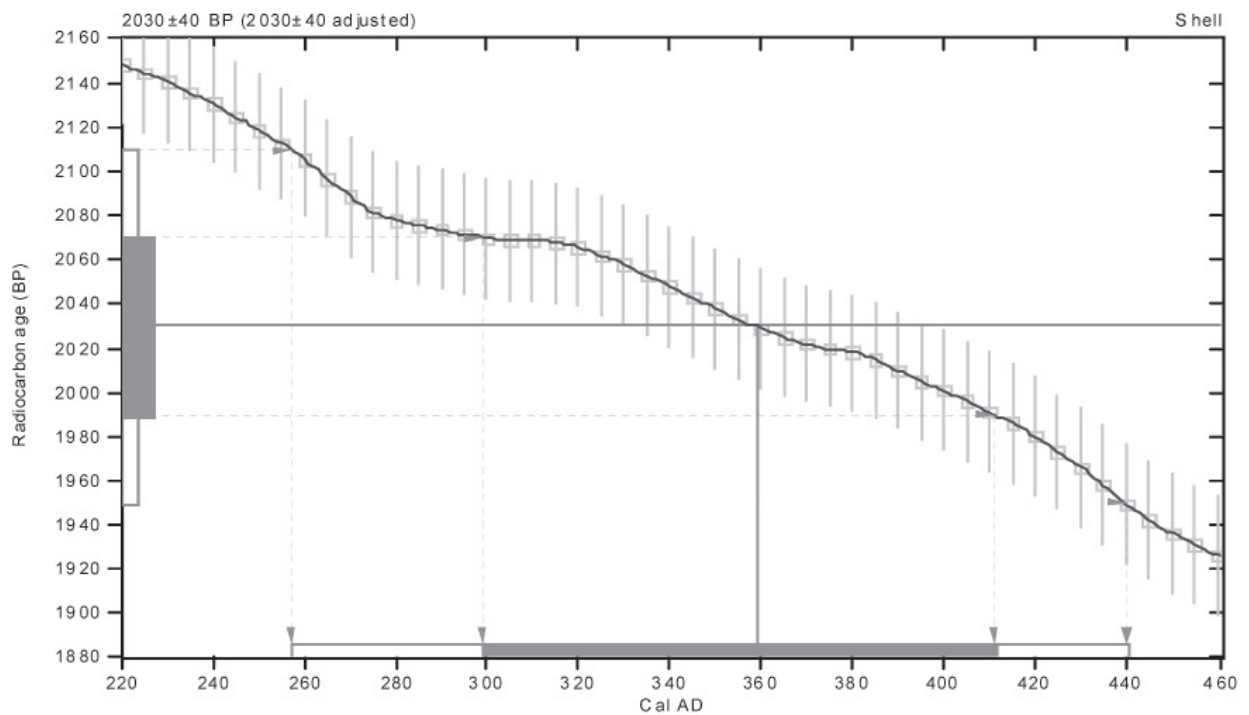
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 260 to 440 (Cal BP 1690 to 1510)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 360 (Cal BP 1590)

1 Sigma calibrated result: Cal AD 300 to 410 (Cal BP 1650 to 1540)
(68% probability)



References:

Database used

MARINE04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.6;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: **Beta-235850**

Conventional radiocarbon age: **2270±40 BP**

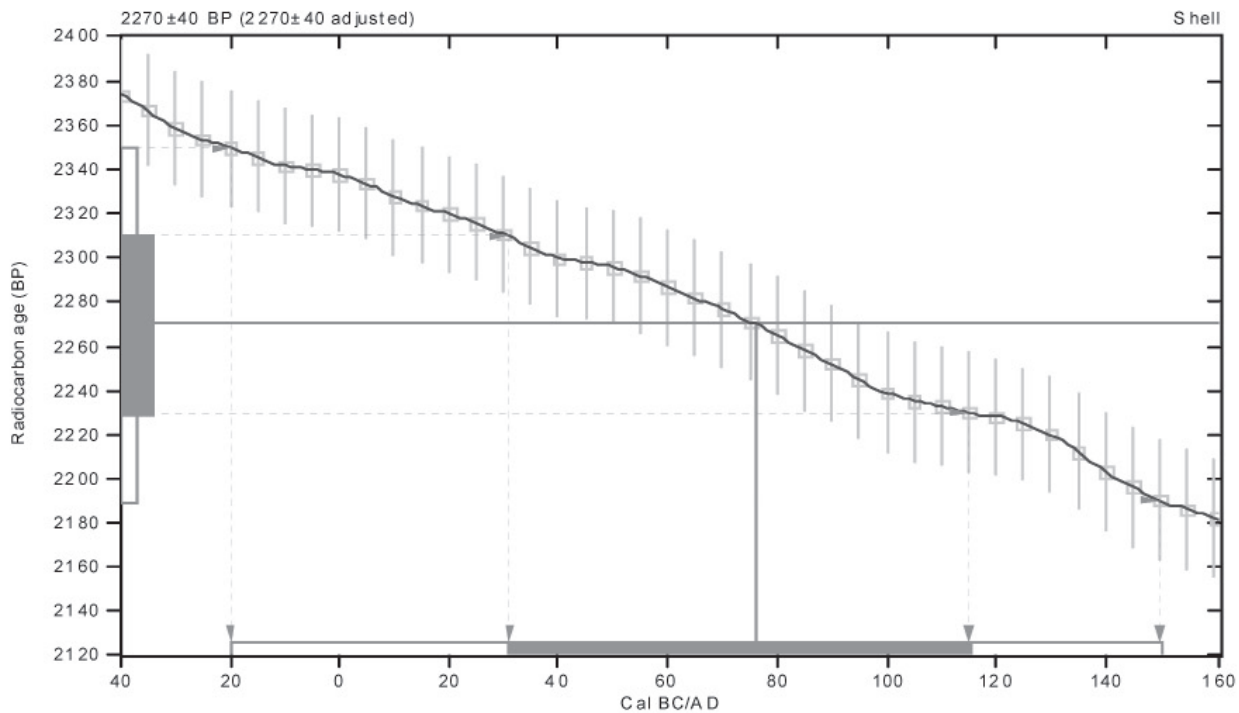
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 20 to Cal AD 150 (Cal BP 1970 to 1800)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal AD 80 (Cal BP 1870)**

1 Sigma calibrated result: Cal AD 30 to 120 (Cal BP 1920 to 1840)
(68% probability)



References:

- Database used*
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.5;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: Beta-235851

Conventional radiocarbon age: 1430±40 BP

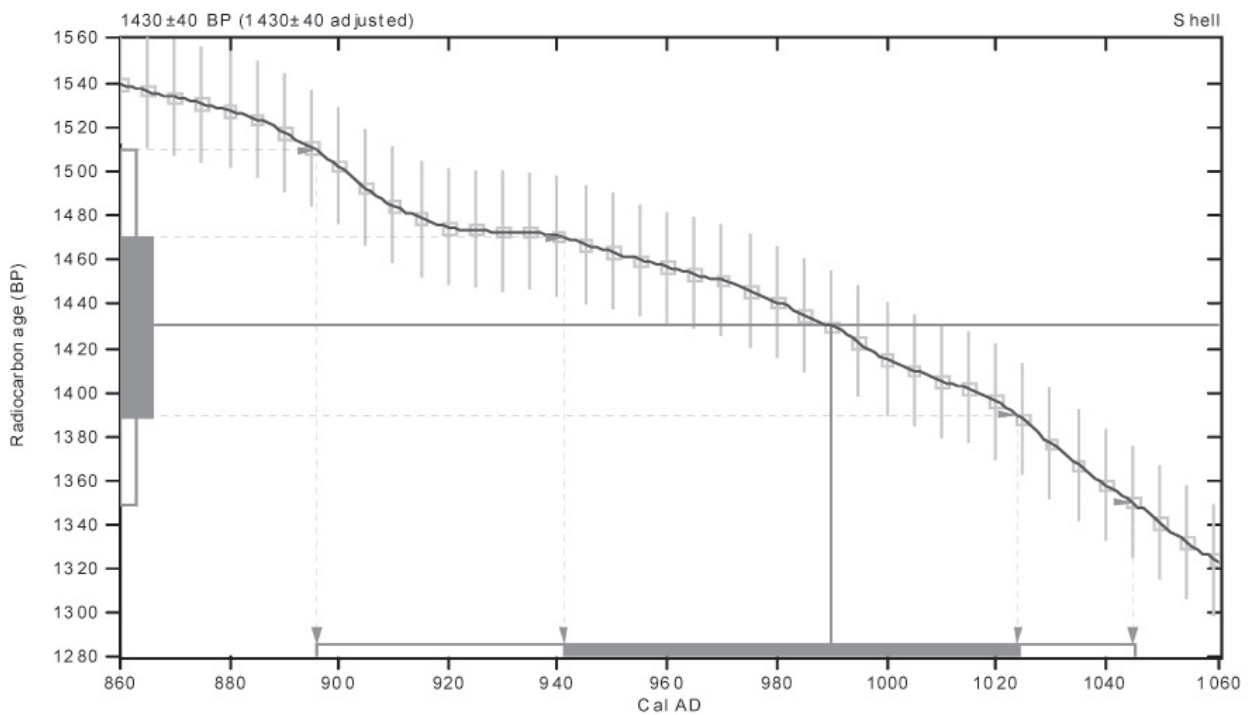
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 900 to 1040 (Cal BP 1050 to 900)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 990 (Cal BP 960)

1 Sigma calibrated result: Cal AD 940 to 1020 (Cal BP 1010 to 930)
(68% probability)



References:

Database used
 MARINE04
Calibration Database
 INTCAL04 Radiocarbon Age Calibration
 IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.8:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235852**

Conventional radiocarbon age: **2310±40 BP**

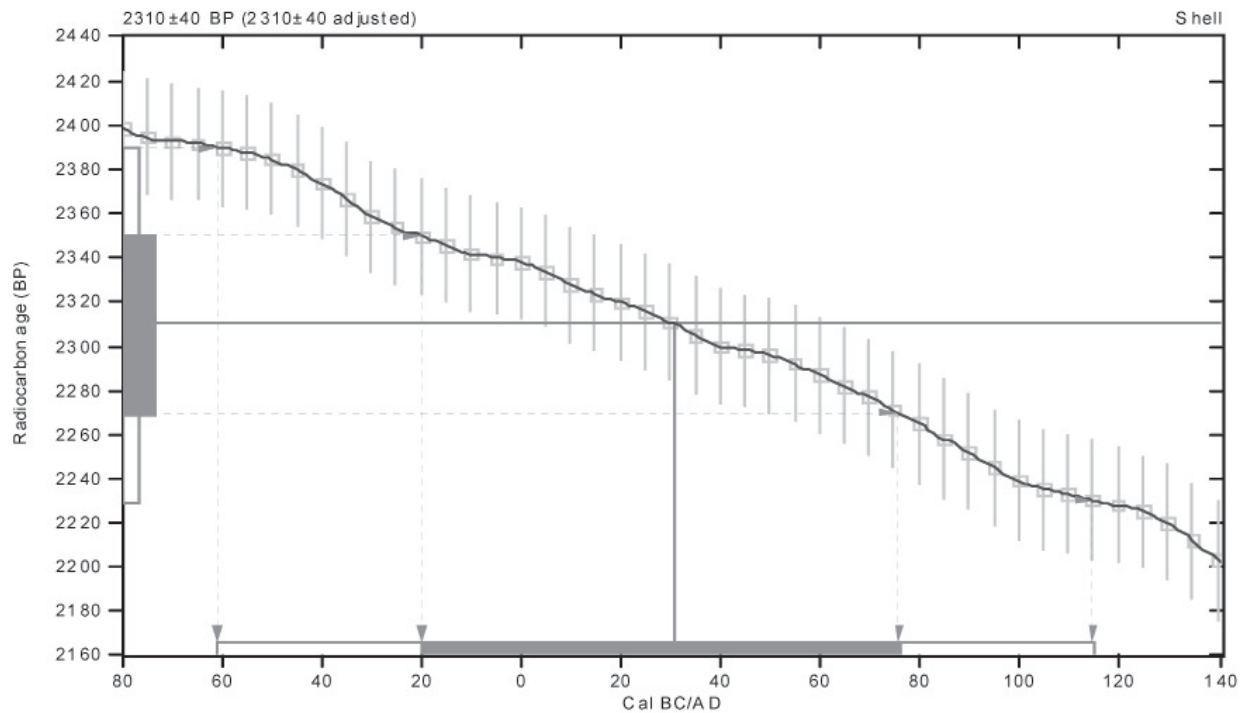
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 60 to Cal AD 120 (Cal BP 2010 to 1840)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 30 (Cal BP 1920)

1 Sigma calibrated result: Cal BC 20 to Cal AD 80 (Cal BP 1970 to 1870)
(68% probability)



References:

- Database used*
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.1;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: **Beta-235853**

Conventional radiocarbon age: **2700±50 BP**

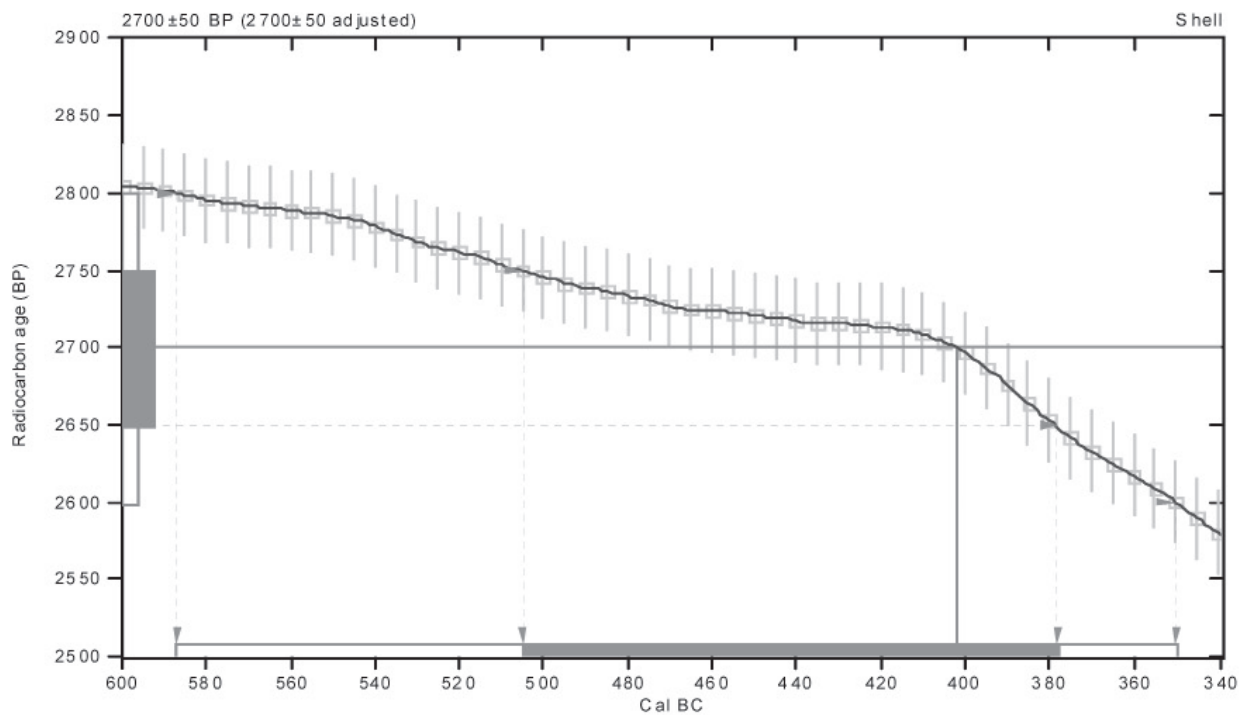
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 590 to 350 (Cal BP 2540 to 2300)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 400 (Cal BP 2350)**

1 Sigma calibrated result: Cal BC 500 to 380 (Cal BP 2460 to 2330)
(68% probability)



References:

Database used

MARINE04

Calibration Database

INTCAL04 Radiocarbon Age Calibration

IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).

Mathematics

A Simplified Approach to Calibrating C14 Dates

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=1.1;Delta-R=0±0;Glob res=-200 to 500;lab. mult=1)

Laboratory number: **Beta-235855**

Conventional radiocarbon age: **1370±40 BP**

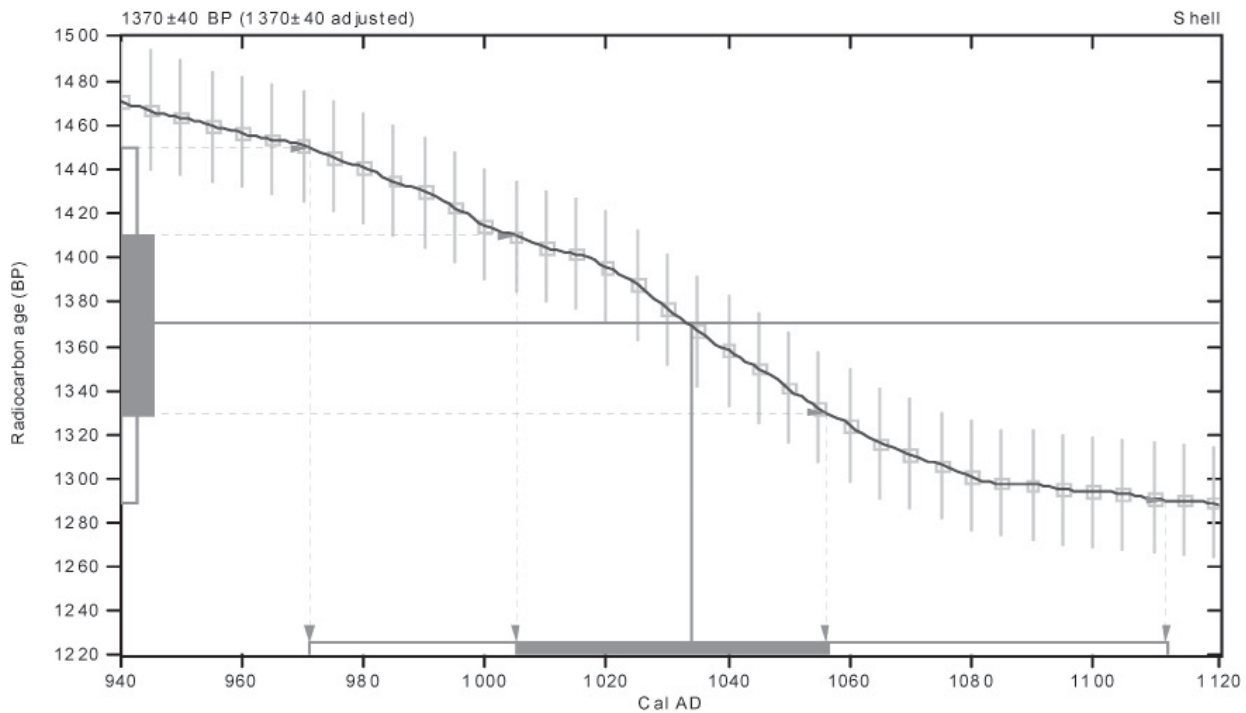
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 970 to 1110 (Cal BP 980 to 840)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 1030 (Cal BP 920)

1 Sigma calibrated result: Cal AD 1000 to 1060 (Cal BP 940 to 890)
(68% probability)



References:

- Database used*
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
Mathematics
A Simplified Approach to Calibrating C14 Dates
Talma, A. S., Vogel, J. C., 1993, Radiocarbon 35(2), p317-322

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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.3:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: Beta-235856

Conventional radiocarbon age: 1510±40 BP

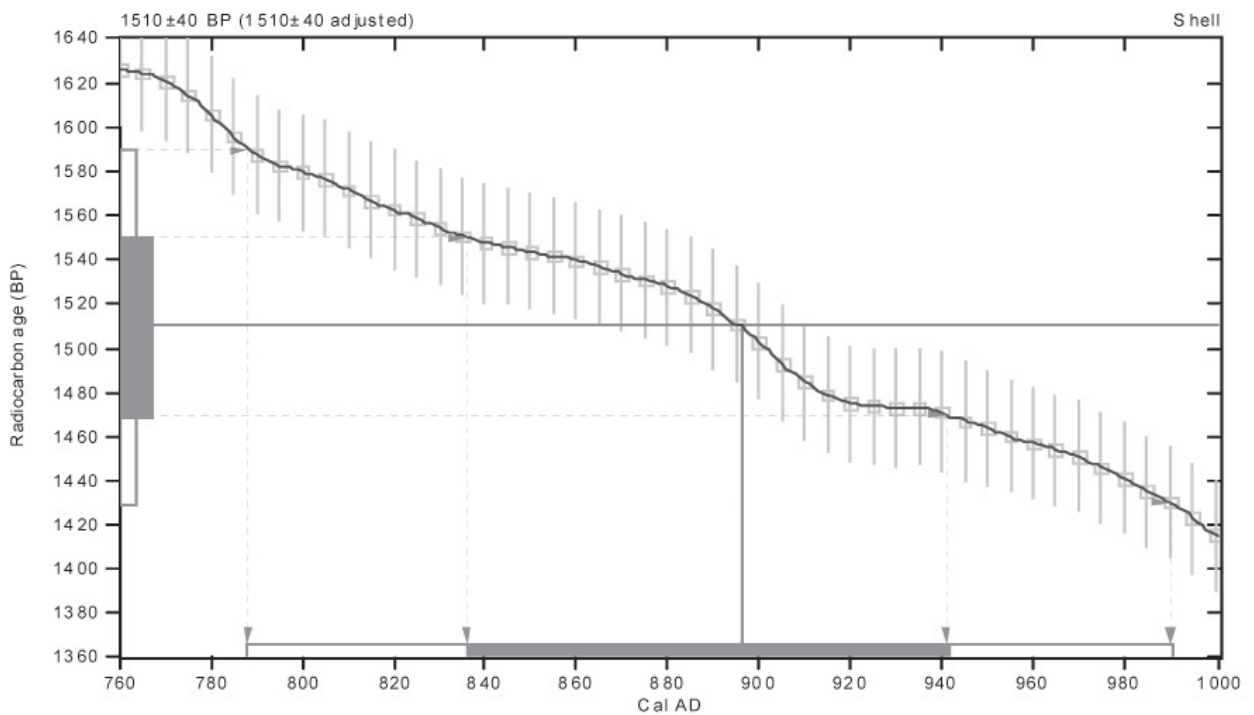
(local reservoir correction not applied)

2 Sigma calibrated result: Cal AD 790 to 990 (Cal BP 1160 to 960)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal AD 900 (Cal BP 1050)

1 Sigma calibrated result: Cal AD 840 to 940 (Cal BP 1110 to 1010)
(68% probability)



References:

Database used
MARINE04
Calibration Database
INTCAL04 Radiocarbon Age Calibration
IntCal04: Calibration Issue of Radiocarbon (Volume 46, nr 3, 2004).
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=0.6:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235857**

Conventional radiocarbon age: **2500±40 BP**

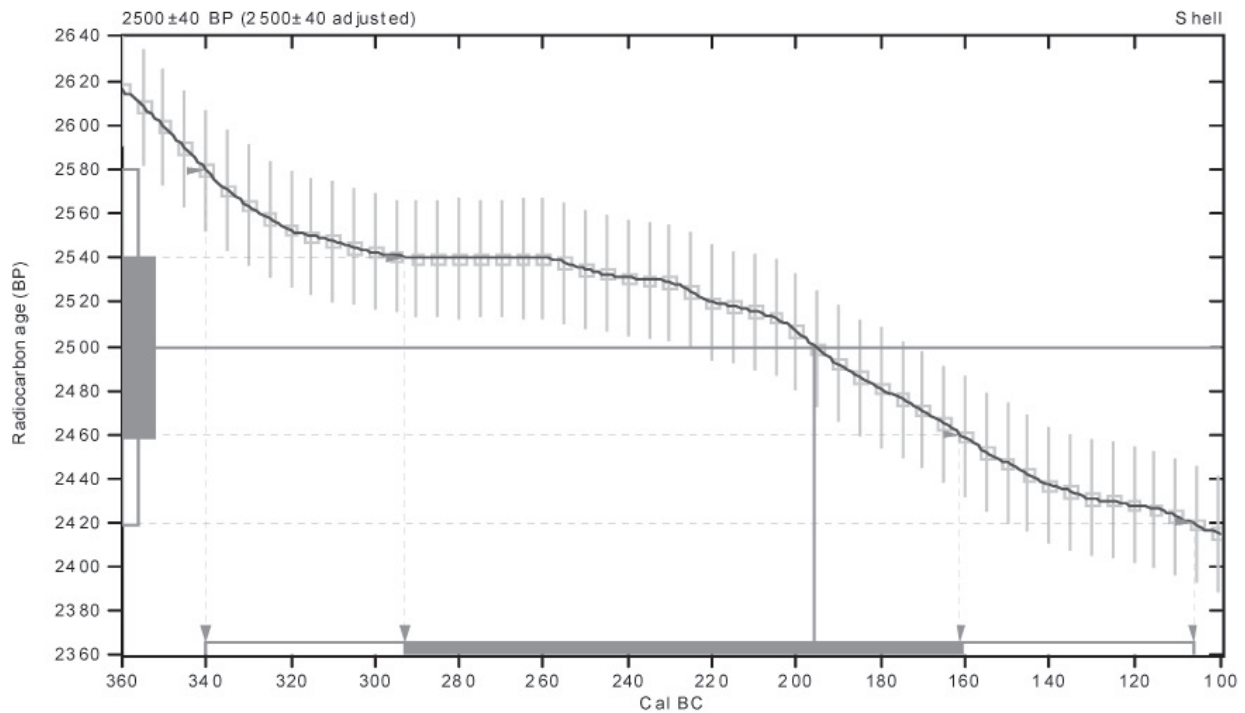
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 340 to 110 (Cal BP 2290 to 2060)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: **Cal BC 200 (Cal BP 2150)**

1 Sigma calibrated result: Cal BC 290 to 160 (Cal BP 2240 to 2110)
(68% probability)



References:

- Database used*
MARINE 04
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INTCAL04 Radiocarbon Age Calibration
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.2:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: Beta-235858

Conventional radiocarbon age: 2480±40 BP

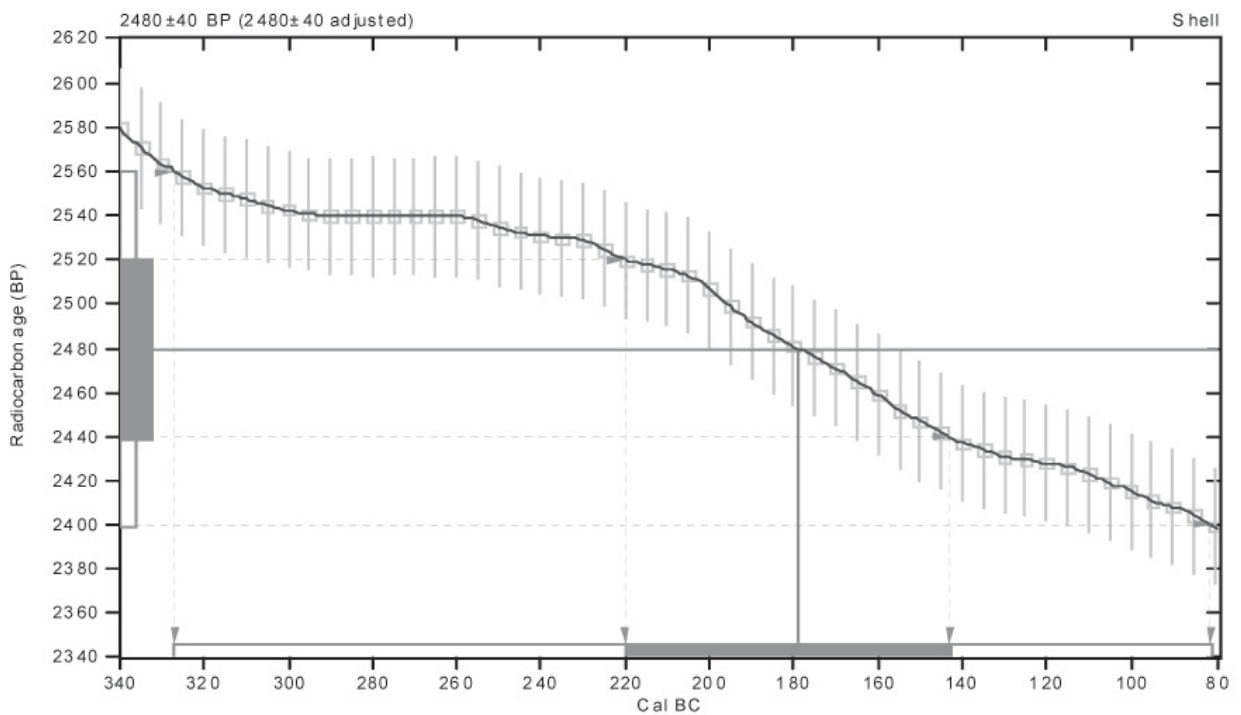
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 330 to 80 (Cal BP 2280 to 2030)
 (95% probability)

Intercept data

Intercept of radiocarbon age
 with calibration curve: Cal BC 180 (Cal BP 2130)

1 Sigma calibrated result: Cal BC 220 to 140 (Cal BP 2170 to 2090)
 (68% probability)



References:

Database used
 MARINE 04
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CALIBRATION OF RADIOCARBON AGE TO CALENDAR YEARS

(Variables: C13/C12=-0.1:Delta-R=0±0:Glob res=-200 to 500:lab. mult=1)

Laboratory number: **Beta-235859**

Conventional radiocarbon age: **3640±40 BP**

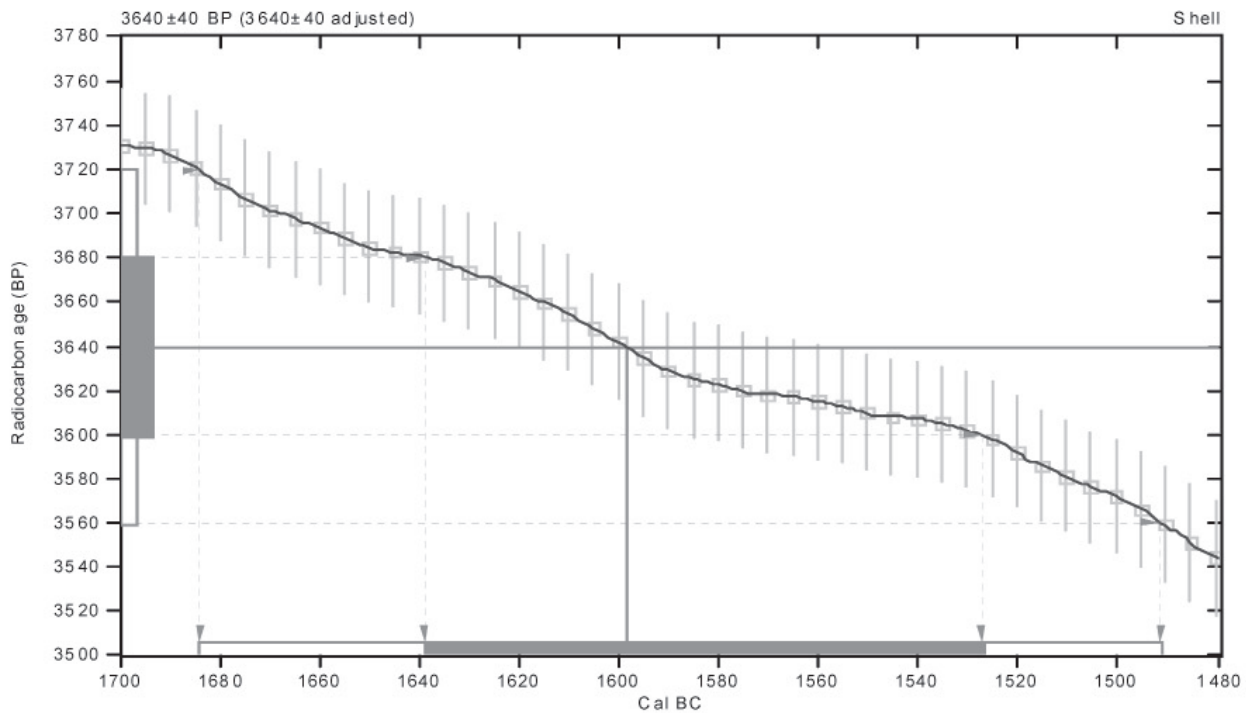
(local reservoir correction not applied)

2 Sigma calibrated result: Cal BC 1680 to 1490 (Cal BP 3630 to 3440)
(95% probability)

Intercept data

Intercept of radiocarbon age
with calibration curve: Cal BC 1600 (Cal BP 3550)

1 Sigma calibrated result: Cal BC 1640 to 1530 (Cal BP 3590 to 3480)
(68% probability)



References:

- Database used*
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