Modeling Coupled Evaporation and Seepage in Ventilated Tunnels

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ABSTRACT

2 Tunnels excavated in unsaturated geological formations are important to activities such 3 as nuclear waste disposal and mining. Such tunnels provide a unique setting for simultaneous 4 occurrence of seepage and evaporation. Previously, inverse numerical modeling of field liquidrelease tests and associated seepage into tunnels were used to provide seepage-related large-scale 5 6 formation properties by ignoring the impact of evaporation. The applicability of such models was 7 limited to the narrow range of ventilation conditions under which the models were calibrated. 8 The objective of this study was to alleviate this limitation by incorporating evaporation into the 9 seepage models. We modeled evaporation as an isothermal vapor diffusion process. The semi-10 physical model accounts for the relative humidity, temperature, and ventilation conditions of the 11 tunnels. The evaporation boundary layer thickness (BLT) over which diffusion occurs was 12 estimated by calibration against free-water evaporation data collected inside the experimental 13 tunnels. The estimated values of BLT were 5 to 7 mm for the open underground tunnels and 20 14 mm for niches closed off by bulkheads. Compared to previous models that neglected the effect of 15 evaporation, this new approach showed significant improvement in capturing seepage 16 fluctuations into open tunnels of low relative humidity. At high relative- humidity values, the 17 effect of evaporation on seepage was very small.

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INTRODUCTION

19 Seepage of liquid water into tunnels is an important phenomenon for subsurface activities 20 such as mining and geologic disposal of nuclear wastes. A key factor affecting the long-term 21 safety of the proposed nuclear waste repository at Yucca Mountain (YM), Nevada, is the seepage 22 of liquid water into waste emplacement tunnels. The rate, chemical composition, and spatial and 23 temporal distributions of seepage are critical factors that determine corrosion of waste canisters, 24 integrity of engineered barriers, and dissolution and mobilization of contaminants and their 25 release to groundwater (Bodvarsson et al., 1999; Finsterle et al., 2003). In unsaturated 26 formations, capillary forces hold the pore water tightly in the formation and prevent it from 27 seeping by gravitational forces into the tunnel – the invisible barrier created by the capillary 28 force is commonly known as a "capillary barrier." Philip and co-workers (Knight et al., 1989; 29 Philip, 1989a; Philip, 1989b; Philip et al., 1989a; Philip et al., 1989b) considered steady-state 30 unsaturated flow around capillary barriers and provided analytical solutions for the critical 31 conditions that trigger seepage into various idealized tunnel goemetries excavated in 32 homogeneous formations. Detailed numerical models have been used to study unsaturated flow 33 in heterogeneous fractured media and seepage into tunnels of various geometries under transient 34 conditions (e.g., Birkholzer et al., 1999; Finsterle, 2000; Finsterle and Trautz, 2001; Li and 35 Tsang, 2003). Site-specific seepage models for the nuclear waste repository at Yucca Mountain, 36 Nevada were developed by calibrating the effective seepage-related parameters against field 37 seepage test data (Finsterle et al., 2003).

Most of the previous numerical models assumed that liquid water leaking into a tunnel drips (seeps) immediately at the place of entry. The potential for evaporation to compete with seepage has been generally ignored, and its effect was lumped with the effective flow parameters 41 of the unsaturated medium (Finsterle et al., 2003). In calibration of the analytical model of Philip 42 et al. (1989b) against field seepage data, Trautz and Wang (2002) accounted for the effect of 43 evaporation by adjusting the field seepage data for evaporation. Because the data were obtained 44 from tests conducted in relatively humid tunnels, the effect of evaporation on the calibrated 45 seepage-related parameter was not significant. However, recent field measurements of seepage 46 and free-water evaporation in ventilated tunnels at Yucca Mountain have shown that seepage rate 47 is significantly influenced by evaporation. The foregoing discussions suggest that the 48 applicability of models that ignore evaporation is limited to similar humidity and temperature 49 conditions under which the calibrations are performed (Finsterle et al., 2003). Such models 50 cannot satisfactorily capture the seepage rate fluctuations when the seepage experiments are 51 conducted under variable humidity and ventilation conditions. More importantly, seepage models 52 that ignore evaporation, and that are calibrated against seepage data under ventilated and/or low 53 humidity conditions are not expected to perform well in predicting future seepage conditions that 54 are expected to be non-ventilated and humid. The preceding observations call for a calibrated 55 seepage model that reliably performs over a wide range of ventilation and humidity conditions.

The objective of this study is to improve the portability of calibrated seepage models by reducing the impact of evaporation on the calibrated effective parameters. Thus, we propose to incorporate evaporation from tunnel walls into the existing seepage models by assuming a firstorder diffusion approximation.

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EVAPORATION IN TUNNELS

Fundamentally, evaporation is a two-step process. The first step involves transition from
liquid to vapor phase at the liquid-vapor interface (vaporization). The second step is the transport
of vapor from the high concentration area at the evaporating surface to the low concentration

area of the ambient air. Accurate modeling of these coupled processes is difficult for several reasons: (1) the first step is a non isothermal phenomenon, and the parameters that govern this process are strongly temperature dependent; (2) the vapor concentration gradient in the boundary layer is strongly influenced by the air flow regime; and (3) the air flow depends on among other things the ambient wind velocity and the roughness of the evaporating surface.

69 Ho (1997) and Or and Ghezzehei (2000) modeled evaporation from individual water 70 droplets attached to tunnel ceilings, assuming constant temperature and humidity conditions. 71 However, the scale of their approach is too small to be incorporated into the larger scale seepage 72 models that represent the discrete dripping process as a continuum flow. Therefore, the 73 evaporation model required in this study should be of an intermediate scale and be compatible 74 with the existing seepage models. The formulation used herein capitalizes on the observed 75 dependence of evaporation rate on tunnel humidity and ventilation conditions, and the 76 availability of high resolution time-series data of relative humidity, temperature and free-water 77 evaporation rate (Trautz and Wang, 2002).

In the following subsections, we introduce an isothermal vapor diffusion model of evaporation and define the problem domain and boundary conditions. This is followed by estimation of the evaporation model parameters, using free-water evaporation data. Finally, a remark on evaporation from porous surface is provided.

82

Isothermal Vapor Diffusion Model

To simplify the first step of evaporation (vaporization) we assume the following: (1) the absorption of latent heat and its effect on the physical properties of the liquid-vapor interface are ignored; (2) the time dependence of the vaporization process (e.g., Zhang and Wang, 2002; Zhang et al., 2001) is neglected; and (3) the vapor partial pressure of the interfacial air is assumed to be under thermodynamic equilibrium. At equilibrium, the air above a flat surface of pure water is considered saturated with vapor; its vapor pressure is denoted by p_s . This saturation vapor pressure rises with temperature. In the temperature range of -10° C to 50°C, the saturation vapor pressure is related to interfacial temperature by (Murray, 1966):

91
$$\ln p_{\rm s} = a \frac{T}{T+b} + c$$
 [1]

92 where a = 21.87, b = 265.5 °C and c = 6.41 are constants, and *T* is the interfacial temperature. 93 For non-flat interfaces (such as capillary menisci) the actual interfacial vapor pressure *p* is 94 related to the interfacial capillary potential by the classic Kelvin equation,

95
$$\ln\left(\frac{p}{p_{\rm s}}\right) = P_C \cdot \frac{M_W}{\rho_W R T}$$
[2]

96 where P_C is the capillary pressure, ρ_W and M_W are the density and molecular mass of liquid 97 water, respectively, and *R* is the universal gas constant. Note that the relative humidity of air is 98 defined as the ratio of the actual partial pressure (*p*) to the saturated vapor pressure (*p_s*)

 $h = p/p_{\rm s}$ ^[3]

100 The second step of evaporation, vapor removal from the interface, is modeled as a first-101 order phenomena described by Fickian diffusion (Rohsenow and Choi, 1961). In one dimension 102 and under constant temperature, the vapor flux (J_y) is given by

$$J_V|_T = -D_V \cdot \frac{\mathrm{d}C}{\mathrm{d}z}$$
[4]

104 where D_V is the vapor diffusion coefficient, which is related to the ambient air pressure (*P*) and 105 air temperature (*T*) by

106
$$D_V = 2.13 \times 10^{-5} \frac{10^5}{P} \left(\frac{T}{273.15}\right)^{1.8}$$
[5]

107 and the vapor concentration C is related to vapor pressure by

108
$$C = \frac{M_W}{RT} \cdot p$$
 [6]

In the subsequent subsection, we define the problem domain and develop the appropriateboundary conditions needed to solve the vapor diffusion equation [4].

111

Velocity and Concentration Boundary Layers

112 In admitting diffusive flux as the primary mechanism for vapor removal from the 113 evaporating surface, we tacitly assume that airflow above the evaporating surface is fully 114 developed and laminar, as illustrated in Fig. 1a. The free-stream air velocity (V^{∞}) is retarded in 115 the vicinity of the evaporating surface because of frictional resistance. The air velocity parallel to the evaporating surface increases from V = 0 at z = 0 (no-slip) asymptotically to $V = V^{\infty}$ at a 116 117 distance sufficiently far away from the surface. For fully laminar flow conditions, the thickness of the boundary layer of retarded velocity (defined as $V \le 0.99V^{\infty}$) is inversely proportional to 118 119 the square root of the free-stream velocity (Rohsenow and Choi, 1961):

120
$$\delta_V \propto 1/\sqrt{V^{\infty}}$$
 [7]

Because the equations that describe laminar air flow parallel to a flat surface and diffusion from a flat surface are analogous (Rohsenow and Choi, 1961), a similar notion of concentration boundary layer holds near the evaporating surface. The vapor concentration profile is illustrated in Fig. 1b. The vapor concentration decreases from an equilibrium value ($C = C^0$) at z = 0 to a value determined by the free-stream humidity at sufficiently far distance. The 126 concentration boundary layer thickness (δ_C) is related to the velocity boundary layer thickness 127 by the Schmidt number,

128
$$Sc = \frac{\delta_C}{\delta_V} = \frac{\mu_a}{\rho_a \cdot D}$$
[8]

where μ_a and ρ_a are the viscosity and density of air, respectively. At 20 °C and 1 atm pressure, the Schmidt number is approximately unity. In the remainder of this paper the subscripts in the boundary layer thickness are dropped and $\delta = \delta_V = \delta_C$. It is evident from [7] and [8] that the concentration BLT (δ) is inversely related to the square root of the free-stream velocity (V^{∞}) and can serve as a direct measure of the tunnel ventilation condition. In a subsequent subsection, estimation of the BLT will provide further elaboration on the dependence of δ on ventilation conditions.

Fig. 1. Schematic description of (a) air velocity and (b) vapor concentration profiles above a
free water surface

138

Boundary Conditions

The domain of the vapor diffusion equation [4] is the concentration boundary layer introduced in the preceding subsection. The boundary condition on [4] corresponding to the equilibrium vapor concentration at the evaporating surface (z = 0) is given by (using [2] and [6]):

143
$$C = C^{0} = \frac{M_{W}}{R T} p_{s} \exp\left[\frac{P_{C}}{\rho_{W}} \frac{M_{W}}{R T}\right]$$
[9]

144 The second boundary condition is at the border of the concentration boundary layer $z = \delta$, where 145 the vapor concentration is defined by the relative humidity (*h*) of the ambient air:

146
$$C = C^{\infty} = \frac{M_W}{R T} p_s h$$
[10]

147 If the boundary conditions change slowly, the evaporation rate can be considered to be at 148 steady state and the concentration gradient dC/dz is constant throughout the boundary layer. 149 Then, the steady state vapor diffusion equation [4] under isothermal conditions is simplified to

$$J_V = -D_v \cdot \frac{C^0 - C^\infty}{\delta}$$
[11]

151 Note that the ratio δ/D_v is commonly referred to as aerodynamic resistance. The isothermal 152 vapor diffusion equation [11] is considered valid for modeling evaporation from tunnel surfaces 153 and free water. Fujimaki and Inoue (2003) found [11] (also known as the bulk transfer equation) 154 to be valid in laboratory evaporation experiments in which the ambient air velocity was on the 155 order of 1 m/s. All the variables of this model are directly related to physical conditions in the 156 tunnel, and all of them, except δ , can be independently determined from measured quantities. 157 The boundary-layer thickness (δ) can be estimated by calibrating [11] against free water 158 evaporation data, as discussed in the next subsection.

159

Estimation of the Boundary-Layer Thickness

Apart from the capillary pressure at the evaporating surface, evaporation from free water and that from a wet porous surface are thus far assumed to be identical processes. Therefore, a controlled evaporation experiment from a still water surface can be used to estimate the vapor concentration boundary layer thickness, which is also applicable to evaporation from wet tunnel surfaces at similar ventilation conditions. Upon substitution of [1], [5], [9], and [10] in [11], and noting that the capillary pressure of the free water surface is $P_C = 0$, we arrive at a free-water evaporation equation,

167
$$J_V = -2.13 \times 10^{-5} \frac{10^5}{P} \left(\frac{T}{273.15}\right)^{1.8} \cdot \left(a\frac{T}{T+b} + c\right) \cdot \frac{M_w}{RT} \frac{1-h}{\delta}$$
[12]

168 According to the isothermal assumption, T denotes the temperature of the evaporating surface 169 and the surrounding air. Assuming the change in conditions that affect evaporation rate is slow 170 compared to the time it takes to reach steady-state evaporation, [12] can be fitted to time-series 171 data of evaporation rate data, measured at known temperature, pressure, and relative humidity 172 conditions. The best-fit δ represents the boundary-layer thickness at the prevailing ventilation 173 condition. However, it should also be noted that uncertainties associated with the assumed simplifications (including isothermal conditions, flat evaporating surface, and laminar airflow) 174 175 are lumped in this parameter. Thus, the boundary-layer thickness should be considered an 176 effective parameter.

177

Evaporation from Porous Surface

178 The surface of an unsaturated porous medium typically consists of solid (matrix of the 179 medium) and pore/fracture (liquid and gas) components, rendering the evaporating surface 180 heterogeneous with respect to vapor concentration, as illustrated in Fig. 2a. During seepage, 181 however, tunnel ceilings are usually covered with liquid films (e.g., Trautz and Wang, 2001), and 182 the vapor concentration could be considered locally homogeneous. For simplicity, we extend 183 this assumption of locally uniform distribution of vapor concentration to the entire tunnel Fig. 184 2b. The vapor concentration at any given location on the tunnel is assumed to be at capillary 185 equilibrium with the pores and fractures of the porous medium. The datum z = 0 for the vapor 186 diffusion is set on the surface of the tunnel (as illustrated in Fig. 2b). Although this assumption is 187 likely to fail at very low saturations (when the liquid is scattered in a few fine pores and

- 188 fractures) it is expected to be of marginal consequence because the evaporation rate under such
- 189 conditions is very low.

190	Fig. 2.	Evaporating surface area of a porous medium: (a) partitioning of the surface into non-
191		evaporating solid and evaporating pores; (b) proposed approach of uniform gas-phase
192		surface. The dark shade denotes vapor in pores and/or fractures.

193COUPLED SEEPAGE AND EVAPORATION

194 In a tunnel constructed in unsaturated formations, the flow velocity of water in the rock is 195 usually stagnated near the crown, resulting in elevated moisture (Philip et al., 1989b). Unlike 196 evaporation from ground surface, where infiltration opposes the evaporation flux, the condition 197 in tunnels is favorable for simultaneous occurrence of evaporation and seepage. Field tests that 198 exhibit simultaneous evaporation and seepage are described below. After field test descriptions, 199 we present a brief description of seepage modeling using the numerical simulators TOUGH2 200 (Pruess et al., 1999) and iTOUGH2 (Finsterle, 1999) and discuss implementation of evaporation 201 in these models.

202

Field Tests

The data reported in this paper were obtained from field tests and measurements conducted at the proposed nuclear waste repository at Yucca Mountain currently under investigation by the US Department of Energy (DOE). Air-injection tests were conducted to characterize the permeability and small-scale heterogeneities of the formation, and liquid-release tests were performed to study seepage phenomena. Relative humidity, temperature, and freewater evaporation were monitored at the test site to assess the evaporation conditions. Detailed description of the site and tests conducted at the site are provided elsewhere (Birkholzer et al., 210 1999; Bodvarsson et al., 1999; Finsterle and Trautz, 2001; Finsterle et al., 2003; Trautz and 211 Wang, 2001; Trautz and Wang, 2002; Wang et al., 1999). This study is concerned with the lower 212 lithophysal welded tuff unit at Yucca Mountain, in which about 80% of the proposed repository 213 is expected to reside. This unit contains many small fractures (less than 1 m long) and is 214 interspersed with numerous lithophysal cavities (0.15 m–1 m in diameter).

215 In the lower lithophysal unit, an 800-m long drift (5 m in diameter) for enhanced 216 characterization of the repository block (ECRB) was excavated off the main Exploratory Studies 217 Facility (ESF) tunnel. Liquid-release and air-injection tests were systematically conducted in this 218 ECRB Cross Drift along boreholes drilled into the ceiling of the Cross Drift at regular intervals. 219 Similar tests were conducted in a short (approximately 15 m long) drift excavated off the Cross 220 Drift (niches). Schematic alignment of the tunnels is shown in Fig. 3a. This paper is concerned 221 with tests conducted at a Cross Drift borehole designated as LA#2 (Fig. 3b) and a short drift 222 known as Niche 5 (Fig. 3c). The tests and measurements conducted in the Cross Drift and Niche 223 5 are briefly described below.

224 Air-injection tests

225 The purpose of the air-injection tests was to estimate absolute permeability of the 226 formation as a basis for the stochastic generation of heterogeneous permeability fields. Short 227 sections of the boreholes (0.3 m in Niche 5, 1.8 m in Cross Drift) were isolated using an 228 inflatable packer system, and compressed air was injected. Air injection was terminated when 229 steady-state pressure was reached. Air-permeability values were derived from the steady-state 230 pressure data according to an analytical solution of LeCain (1995). Permeabilities determined 231 from air-injection tests were considered representative of the absolute permeability of the test 232 interval.

233 Liquid-release Tests

234 Liquid release tests were conducted in boreholes drilled above tunnels to evaluate 235 seepage into waste emplacement drifts. The alignment of the boreholes and test intervals are 236 schematically shown in Fig. 3. The liquid release boreholes in the Cross Drift were 237 approximately 20 m long, drilled into the ceiling of the Cross Drift at a nominal inclination of 238 15° from the horizontal. Liquid release data from a borehole designated as LA#2 were used in 239 this study. The borehole was partitioned into three zones (designated as Zone 1, Zone 2, and 240 Zone 3) available for liquid release testing. The distances from the middle of the liquid-release 241 zones to the drift crown were 1.58 m, 2.84 m, and 4.10 m for Zone 1, Zone 2, and Zone 3, 242 respectively. The liquid release boreholes in Niche 5 were near horizontal. Of the six boreholes 243 available for tests, data from boreholes #4 and #5 were used in this study. The liquid release tests 244 were performed by injecting water into a test interval isolated by inflated rubber packers. Water 245 that seeped into the tunnels was captured and measured using automated recording devices.

246

Relative Humidity and Temperature Measurements

The Cross Drift was actively ventilated during regular working hours, thus the relative humidity of the tunnel was usually low. To mitigate the effect of evaporation in the seepage process, the seepage collection interval was guarded using curtains on both ends. Because Niche 5 was isolated from the actively ventilated Cross Drift by a bulkhead, the relative humidity was relatively high. To aid in the estimation of evaporation during the liquid release tests, the relative humidity and temperature of the air inside and outside of the curtains (for the Cross Drift) and in front of and behind the bulkhead (for Niche 5) were monitored.

254	The evaporation rate from still water was measured by monitoring the level (mass) of
255	water in evaporation pans placed within the space enclosed by the seepage capture tray and end
256	curtains (for the Cross Drift tests) and behind the bulkhead (for Niche 5).

257	Fig. 3.	3. Schematic alignment of tunnels and boreholes: (a) parts of the Exploratory Studies	
258		Facility (ESF) tunnel and Enhanced Characterization of the Repository Block (ECRB)	
259		cross-drift; (b) liquid release test setup in the Cross Drift, including liquid release	
260		intervals and liquid injection and seepage collection equipment; and (c) vertical section	
261		of Niche 5 along with location of all the test boreholes.	

TOUGH2/iTOUGH2 Seepage Model

A detailed description of the numerical models developed for flow in fractured formation around a tunnel and associated seepage into the tunnel using TOUGH2/iTOUGH is given by Finsterle et al (2003). A summary follows.

The TOUGH2 code is an integral finite difference simulator that represents unsaturated flow at the scale of individual grids by Richards' equation (Bear, 1972; Pruess et al., 1999)

268
$$\phi \rho \frac{\partial}{\partial t} S_e = \operatorname{div} \left[k \frac{\rho}{\mu} \nabla (P_C + \rho g z) \right]$$
[13]

The appropriateness of using this continuum approach to simulate water flow through unsaturated fractured rock was shown by Finsterle (2000). The effective permeability (k) and capillary pressure (P_C) are functions of liquid saturation as given by van Genuchten's models (1980)

273
$$k = k_a S_e^{1/2} \left[1 - \left(1 - S_e^{1/m} \right)^m \right]^2$$
[14]

274
$$P_C = -\frac{1}{\alpha} \left[S_e^{-1/m} - 1 \right]^{1-m}$$
[15]

where k_a is the absolute permeability, $1/\alpha$ and *m* are fitting parameters with $\alpha > 0$ and 275 0 < m < 1, and the effective saturation, S_e , is defined as $S_e = (S - S_{lr})/(1 - S_{lr})$, with S_{lr} being the 276 residual liquid saturation. While the k_a values were considered spatially heterogeneous, the $1/\alpha$, 277 m, and S_{lr} parameters were summed to be homogeneous for a given test bed (Finsterle et al., 278 2003). The absolute permeability, k_a , was derived from the air-injection tests. The van 279 Genuchten *m* parameter and the residual saturation were taken to be m = 0.608 and $S_{lr} = 0.01$, 280 281 respectively (Finsterle et al., 2003). The van Genuchten capillary strength parameter $1/\alpha$ was 282 estimated through inverse modeling. In the numerical seepage model, the condition for seepage 283 is determined by the total water-potential gradient at the connection between the porous medium 284 and the tunnel, as depicted in Fig. 4. The flow rate along the connection between the porous 285 medium and the tunnel is given by

286
$$q_z = k \frac{\rho}{\mu} \frac{\Delta P + \rho g z}{\Delta z}$$
[16]

where ΔP denotes the capillary pressure difference across the distance between the last formation node and the tunnel node Δz . The nodal distance Δz is chosen to be a representative of the average length of fractures intersecting the tunnel that are not draining laterally (Finsterle et al., 2003). From [16], and assuming that the capillary pressure in the opening is zero, it follows that downward seepage ($q_z > 0$) occurs only when the following condition is satisfied:

 $-P_C^* > \rho g \Delta z \qquad [17]$

293 where P_C^* is the threshold capillary pressure at the last node adjacent to the opening. The critical 294 capillary pressure $P_C^* = -\rho g \Delta z$ depends on the grid size or nodal distance of the numerical

- model. According to [17], the tunnel surface does not need to be fully saturated for seepage to
- commence as in the case of unfractured homogeneous porous media (Philip et al., 1989b).

297	Fig. 4.	Schematic description of the seepage and evaporation connections between nodes that
298		represent the rock of the tunnel wall and the tunnel.

299

Implementation of Evaporation in TOUGH2

300 While seepage occurs only when the critical condition given in [17] is satisfied, vapor 301 flow from/to tunnel walls to/from tunnel air occurs as long as there is vapor pressure 302 disequilibrium between them. Coupling of the seepage and evaporation processes is illustrated in 303 Fig. 4. Mass-transfer rate of water, including seepage, is represented in TOUGH2 by equations 304 similar to [16], where the driving force is pressure gradient. To incorporate evaporation into the 305 existing model without significant changes to the governing flow equations, we must rewrite the 306 concentration-gradient dependent diffusion equation [11] in the form of equation [16]. Noting 307 that the connection length Δz denotes the vapor concentration boundary layer thickness δ , the 308 equivalent evaporative permeability can be written as

309
$$k_{eq} = D_{\nu} \frac{\mu}{\rho} \left(\frac{C^0 - C^{\infty}}{P_C^0 - P_C^{\infty}} \right)$$
[18]

where the variables with a superscript of 0 correspond to the tunnel wall and those with a superscript of ∞ denote the tunnel air. The capillary pressure of the tunnel P_C^{∞} is equivalent to the relative humidity [3] of the tunnel, as described by Kelvin's equation [2]. The vapor concentrations are computed according to [11] and [12]. Equation [18] was implemented in TOUGH2 as a special evaporation connection. When the conditions for both evaporation and 315 seepage permit, the total mass flow from the tunnel wall to the tunnel is considered as the sum of 316 both.

317

Numerical Meshes

318 Different numerical models were constructed to simulate liquid-release tests and seepage 319 into the underground openings at different test locations. Three-dimensional meshes of the test 320 sites were generated with grid sizes of 0.3 m \times 0.1 m \times 0.1 m for the Cross Drift and 0.1 m \times 321 $0.1 \text{ m} \times 0.1 \text{ m}$ for Niche 5 (see Fig. 5). For the Cross Drift meshes, a circular cylindrical tunnel 322 of 5 m diameter was removed from the center of the mesh to represent the tunnel. Only one half 323 of the symmetric mesh was used in the simulations to save computational load. For the Niche 5 324 meshes, surveyed niche geometry was removed from the numerical mesh to replicate the test 325 sites. The liquid-release boreholes are indicated in Fig. 5 by bold black lines, and the white 326 sections at the middle of the boreholes represent the injection intervals. The Cross Drift borehole 327 is inclined while the Niche 5 boreholes are parallel to the centerline of the niche. The Cross Drift 328 mesh in Fig. 5a represents the Zone 2 test interval. In Fig. 5b and Fig. 5c, boreholes #4 and #5 329 are revealed, respectively (see also Fig. 3c). Notice that the injection intervals in boreholes #4 330 and #5 are located at 3–3.5 m and 8.5–8.8 m, respectively, from the borehole collars; hence, the 331 respective tunnel outlines are different.

Fig. 5. Numerical meshes of (a) Niche 5 with borehole #4, (b) Niche 5 with borehole #5, and
(c) the Cross Drift, along with a typical realization of the correlated stochastic
permeability field. Bold black lines denote the liquid-release boreholes, and the white
section in the middle of the boreholes is the injection interval.

336 The spatial structure of the Niche 5 permeability data was analyzed using the GSLIB 337 module GAMV3 (Deutsch and Journel, 1992) and a spherical semivariogram was fitted to the 338 resulting variogram. Because only six permeability data were available from the Cross Drift, 339 assumed spherical variogram parameters were used. Recall that the permeability of the Cross 340 Drift was measured on 1.8 m long intervals of the boreholes, and the standard deviation of the 341 measured data was 0.21. The variability of the permeability on the scale of the 0.3 m long 342 gridblock was expected to be greater than the measurement interval. For the purpose of 343 generating a heterogeneous field, the permeability was taken to be log-normally distributed with 344 a variance (sill) value of 1 order of magnitude. Computed and prescribed geostatistical 345 parameters (Table 1) were used to generate spatially correlated permeability fields, using the 346 sequential indicator simulation (SISIM) module of the GSLIB (Deutsch and Journel, 1992). 347 Multiple realizations of the permeability field were generated and mapped to the numerical 348 meshes. Representative permeability field realizations for the Cross Drift and Niche 5 are shown 349 in Fig. 5.

350 351

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Table 1. Mean, standard deviation, and correlation length of log-permeability data collected in the Cross Drift and Niche 5. The values in parentheses are prescribed values because the number of measurements was not adequate to compute the respective parameters.

		Mean log (k) [m²]	Std. Dev. [m ²]	Spherical Variogram		
Location	n			Sill Value [log(k) ²]	Correlation length [m]	Nugget effect [log(k) ²]
Niche 5	61	-10.95	1.31	1.81	0.91	0.02
Cross Drift	6	-10.73	0.21	1.0	0.2	-

353

The tunnels were represented in the seepage models by two types of overlapping gridblocks, one corresponding to seepage and the other to evaporation. The seepage gridblocks were assigned a zero capillary pressure, whereas the evaporation gridblocks were assigned a capillary pressure and vapor concentration corresponding to the tunnel relative humidity of the tunnel, as given by [2] and [3]. No-flow boundary conditions are specified at the left, right, front, and back sides of the model. A free-drainage boundary condition is applied at the bottom to prevent an unphysical capillary boundary effect.

361

RESULTS AND DISCUSSIONS

362 **Evaporation Boundary Layer** The evaporation data collected in Niche 5 were used to calibrate the evaporation model. 363 364 The data were grouped into three classes based on airflow velocity (ventilation): (1) inside Niche 365 5 without ventilation; (2) outside Niche 5 with active ventilation; and (3) outside Niche 5 without 366 active ventilation, the regime usually encountered during nights and weekends. In Fig. 6, the 367 measured relative humidity, and temperature, and evaporation rates from still water are plotted. 368 The evaporation model [12] was fitted to the measured data by adjusting the boundary layer 369 thickness. The best-fit estimates of the boundary layer thickness are listed in Table 2.

Fig. 6. Temperature, humidity, and evaporation rate data, along with model fit of the evaporation data for inside and outside of Niche 5.

In agreement with the theoretical assessment (Equation [7]), the estimated δ showed an inverse relationship with the ventilation conditions. Inside Niche 5, the air was the calmest because it was isolated from the Cross Drift by a bulkhead (see Fig. 3). As a result, the thickest boundary layer (20 mm) was obtained inside Niche 5. Fig. 6 shows that the relative humidity 376 outside Niche 5 increases at nights and during weekends when active tunnel ventilation is turned 377 off. However, this increase in relative humidity is insufficient to explain the observed decrease in 378 evaporation. Therefore, as shown in Fig. 6, reduced air ventilation during nights and weekends is 379 also accompanied by an increase in the thickness of the boundary layer. The estimated boundary-380 layer-thickness values and Equation [7] suggest that the air velocity outside Niche 5 is higher 381 than the inside by factors of 7 (without active ventilation) and 16 (with active ventilation). These 382 results confirm the applicability of Equation [12] to describe the effects of humidity, 383 temperature, and ventilation on evaporation rate.

384

Table 2. Summary of estimated boundary layer thickness for Niche 5 and their application.

Location of Experiment	δ (mm)	Used For Simulation of Liquid- Release Tests in
Inside Niche 5	20.0	Niche 5
Outside Niche 5, ventilation off	7.5	Cross Drift (with end curtains)
Outside Niche 5, ventilation on	5.0	Not used

385

386

Coupled Seepage and Evaporation

In this section, simulations of coupled seepage and evaporation are compared with measured seepage rate data. The software iTOUGH2 (Finsterle, 1999) was used to match the simulated seepage rate with the measured values by adjusting the free capillary strength parameter $(1/\alpha)$ (Finsterle et al., 2003). The corresponding evaporation rate from the tunnel walls simulated using the tunnel relative humidity and calibrated boundary layer thickness. **392** Niche 5

393 Here, two different data sets from liquid release tests conducted in boreholes #4 (October, 394 2002) and #5 (July 2002) are compared with the Niche 5 seepage models. The liquid release rate, 395 seepage rate, and relative humidity data as well as modeled liquid release rate and fitted seepage 396 rate are shown in Fig. 7. The best-fit $1/\alpha$ values were 671 ± 223 Pa and 740 ± 339 Pa for 397 boreholes #4 (30 inversions) and #5 (24 inversions), respectively. The measured seepage rates 398 attained a steady-state flow rate after several days. Because the early-time transient data are 399 biased by storage (e.g., in lithophysal cavities and matrix) and/or fast flow paths connecting the 400 injection interval to the tunnel ceilings, the model was fitted to the late-time steady state data. In 401 the simulations, the relative humidity was kept constant at 0.85 to match with the lowest steady-402 conditions observed during the borehole #4 tests.

403 Fig. 7. Calibration of seepage-rate data from liquid-release tests conducted in Niche 5.
404 Calculated seepage rate curves show only one of the multiple inversions.

405 To quantify the impact of evaporation on seepage over the observed high relative 406 humidity range (0.85-0.99), the calibrated seepage model of borehole #4 was used to simulate 407 seepage and evaporation at relative humidity values of 0.85, 0.95, and 0.99. The resulting steady 408 state seepage and evaporation rates (on Day 10) are plotted as percentages of the liquid release 409 rate in Fig. 8. At a relative humidity of 0.85, the evaporation rate from the entire niche wall 410 surface and the seepage rate are comparable in magnitude. As the relative humidity was 411 increased, the steady-state evaporation rate showed a drastic decrease, while the corresponding 412 seepage rate increased only slightly. Thus, at these high relative humidity conditions, the main 413 impact of evaporation is on the quantity of liquid diverted around the tunnel.

414 Fig. 8. Effect of high relative humidity on evaporation and seepage rates.

415 ECRB Cross Drift

416 In this subsection, two different data sets from liquid release tests conducted in borehole 417 LA2, Zone 2 and Zone 3, are compared with the ECRB Cross Drift seepage model. The liquid 418 release rate, seepage rate, and relative humidity data, as well as modeled liquid-release rates and 419 fitted seepage rates, are plotted in Fig. 9. The best-fit capillary-strength parameter $1/\alpha$ were 557 420 \pm 56 Pa for zone 2 and 535 \pm 58 Pa for zone 3, based on 21 and 19 inversions, respectively. Note 421 that both of the liquid-release tests were conducted concurrently. The measured and simulated 422 seepage rate fluctuations were strongly correlated to the drastic changes in relative humidity 423 (hence, evaporation). The model captured this evaporation effect satisfactorily, tracking 424 increases in measured seepage rates as relative humidity increases and vice versa.

425 Fig. 9. Calibration of seepage-rate data from liquid-release tests conducted in the ECRB Cross 426 Drift. Calculated seepage rate curves show only one of the multiple inversions.

427 The interplay between relative humidity fluctuation and dynamics of flow and ceiling 428 wetness at different times during the test in Zone 2 are visualized in Fig. 10. During this test, the 429 liquid release rate was relatively stable (steadily increasing from 31 mL/min on Day 0 to 34 430 mL/min on Day 34). However, the relative humidity fluctuated between 30% and 90% during 431 this period. Fig. 10 shows snapshots of the liquid saturation distribution on Days 0, 10, 20, and 432 30. Just before the test began, the drift wall has dried out because of the low relative humidity in 433 the drift. The liquid saturation at this time was in equilibrium was the assumed background 434 percolation flux of 2 mm/yr. On day 10 day of injection (relative humidity \sim 70%), water 435 reached the crown of the drift, seepage has started, water was being diverted around the drift, and

436 wet plume has reached approximately to the elevation of the spring line. After 20 days, however, 437 the plume has shrunk significantly because of reduced humidity (approximately 12%) and 438 increased evaporation. Moreover, the seepage rate and seepage locations (indicated by inverted 439 triangles) have decreased. Before the 30-day time mark, the relative humidity rose up to 440 approximately 80%; thus, the evaporation rate was reduced, the wet plume grew, and seepage 441 rate and number of seeps increased. In general, despite the high liquid release rate, the flow 442 regime remained unsaturated. The liquid saturation was highest near the drift crown, which 443 induces a capillary pressure gradient that promoted flow diversion around the drift (capillary 444 barrier effect). Seepage and evaporation removed water from the formation as water flows 445 around the drift, limiting the spread of the wetted region on the drift wall.

446	Fig. 10.	Liquid saturation distribution simulated with model calibrated against seepage-rate data
447		from liquid-release tests conducted in the Cross Drift borehole LA#2, Zone 2 at 0, 10,
448		20, and 30 days after the start of the liquid release tests. Note the correlation of tunnel
449		wall wetness to tunnel relative humidity.

450

SUMMARY AND CONCLUSIONS

In this paper, we (1) estimated the evaporative boundary-layer thickness by calibrating a semi-physical evaporation model, which considers isothermal vapor diffusion; (2) calibrated a heterogeneous fracture-continuum model against seepage-rate data; and (3) tested the effect of evaporation on seepage predictions. The major conclusions of this study are listed below:

The simplified vapor-diffusion approach of modeling evaporation was found to be effective
 in capturing the roles of the important environmental conditions that affect evaporation –
 namely, relative humidity, temperature, and ventilation. Calibrated thicknesses of the

evaporation boundary layer were obtained for three ventilation conditions representing theconditions at the liquid-release test sites at Yucca Mountain.

460 2. We found that evaporation reduces seepage significantly in tests conducted under 461 ventilated conditions. Therefore, it is important to account for evaporation effects when 462 calibrating a seepage process model against liquid-release-test data collected under 463 ventilated conditions. In contrast, the impact of evaporation on seepage rate was minimal in 464 closed-off niches, where relative humidity values were generally high. Thus, when using 465 data obtained from closed-off and/or artificially humidified niches, ignoring the effect of 466 evaporation is expected to introduce little error in the estimation of seepage-relevant 467 parameters.

468 3. The classification of ventilation regimes is based on crude assessment of the tunnel 469 environment. Bearing of external wind velocity variations (note that the Cross Drift is 470 connected to the air outside the ESF) was not accounted for. The matching between 471 measured evaporation rate and model predictions can be improved if accurate measurement 472 of air velocity in the tunnels was made.

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542	

Figures



Fig. 1. Schematic description of (a) air velocity and (b) vapor concentration profiles above a free water surface



Fig. 2. Evaporating surface area of a porous medium: (a) partitioning of the surface into nonevaporating solid and evaporating pores; (b) proposed approach of uniform gas-phase surface.



Fig. 3. Schematic alignment of tunnels and boreholes: (a) parts of the Exploratory Studies Facility (ESF) tunnel and Enhanced Characterization of the Repository Block (ECRB) cross-drift; (b) liquid release test setup in the Cross Drift, including liquid release intervals and liquid injection and seepage collection equipment; and (c) vertical section of Niche 5 along with location of all the test boreholes.



Fig. 4. Schematic description of the seepage and evaporation connections between nodes that represent the rock of the tunnel wall and the tunnel.



Fig. 5. Numerical meshes of (a) Niche 5 with borehole #4, (b) Niche 5 with borehole #5, and (c) the Cross Drift, along with a typical realization of the correlated stochastic permeability field. Bold black lines denote the liquid-release boreholes, and the white section in the middle of the boreholes is the injection interval.



Fig. 6. Temperature, humidity, and evaporation rate data, along with model fit of the evaporation data



Fig. 7. Calibration of seepage-rate data from liquid-release tests conducted in Niche 5. Calculated seepage rate curves show only one of the multiple inversions.



Fig. 8. Effect of high relative humidity on evaporation and seepage rates.



Fig. 9. Calibration of seepage-rate data from liquid-release tests conducted in the ECRB Cross Drift. Calculated seepage rate curves show only one of the multiple inversions.



Fig. 10. Liquid saturation distribution simulated with model calibrated against seepage-rate data from liquid-release tests conducted in the Cross Drift borehole LA#2, Zone 2 at 0, 10, 20, and 30 days after the start of the liquid release tests. Note the correlation of tunnel wall wetness to tunnel relative humidity.