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Investigation of Rifting Processes in the Rio Grande Rift Using Data
from Unusually Large Earthquake Swarms

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I. Introduction

San Acacia Swarm in the Rio Grande Rift. Because the Rio Grande rift is one of the best seismically instrumented rift zones in the world, studying its seismicity provides an exceptional opportunity to explore the active tectonic processes within continental rifts. We have been studying earthquake swarms recorded near Socorro in an effort to link seismicity directly to the rifting process. For FY94, our research has focused on the San Acacia swarm (Figure 1), which occurred 25 km north of Socorro, New Mexico, along the accommodation zone between the Albuquerque-Belen and Socorro basins of the central Rio Grande rift. The swarm commenced on 25 February 1983, had a magnitude 4.2 main shock on 2 March and ended on 17 March, 1983.

The San Acacia Swarm and Its Tectonic Setting. The San Acacia swarm was centered near the point of maximum geodetic uplift thought to be due to inflation of the Socorro Magma Body (SMB). The SMB [Hartse et al., 1992] is a greater than 2000 km² sill-like feature at about 19 km depth. The lateral extent of the magma body (Figure 1) has been estimated using reflected phases identified on microearthquake seismograms [Rinehart et al., 1979], and its internal structure has been studied by Ake and Sanford [1988] and Brocher [1981]. Maximum surface uplift rates of 2 mm/yr over the magma body have been reported by Larsen et al. [1986].

Data. The swarm was recorded by a permanent, telemetered network of from 8 to 10 stations (Figure 1). Recording was supplemented with three to five temporary stations occupying 12 sites in the epicentral region for 10 days following the $M_d=4.2$ main shock. Location of hypocenters for swarm events did not depend solely on readings of first-arriving P phases and their associated S phases. Whenever observed, we included reflected phases P_2P , S_2P and S_2S from the mid-crustal magma body in the location process using the joint hypocenter determination program SEISMOS written by Hartse [1991]. The

use of reflected phases has been shown to reduce focal depth errors of Socorro area earthquakes by a factor of three over the use of direct phases only [Hartse, 1991].

II. Progress During the Reporting Period

A. Major Findings

Hypocenters. We used the velocity model estimated by Hartse [1991] for the central Rio Grande Rift in this study, and we estimated station corrections that are specific to the San Acacia swarm. Station corrections (also known as station delays) account for velocity and thickness variations within the uppermost crust, above basement rocks. A positive correction generally indicates thick deposits of low-velocity material (relatively unconsolidated sediments) below a station. A negative or very small correction indicates that the station is sited where basement rocks are at (or very close to) the surface. To test that the station corrections we found are geologically reasonable, we examined gravity data for the San Acacia area. Most of the gravity measurements were made by previous investigators, but we made some measurements to supplement the earlier studies. The Bouguer gravity map (Figures and 3) clearly shows a gravity low associated with the rift. However, along the low, a local gravity high forms a saddle between the Albuquerque-Belen and Socorro basins. We believe that the gravity saddle identifies an area where basement rocks are close to the surface. The smallest station corrections we estimated are from sites that are over this gravity saddle. Thus, the gravity and seismic data are in agreement.

Hypocenters for 204 events in the swarm were computed of which 102 were considered high quality, latitude, longitude and depth errors of ≤ 0.5 km (Figures 3 and 4). For the well-constrained data set, the focal depths are evenly distributed between 3.3 and 6.3 km depth. Because the swarm is located on the saddle between the two basins, the shallowest hypocenters are below the Phanerozoic-Precambrian interface. Therefore the standard model with station corrections is an adequate velocity model for locations and fault mechanism analysis of swarm events.

Cross sections of the earthquake hypocenters roughly define an ellipsoid with the major axis striking ENE and plunging 45° along strike. (D-D' and H-H' in Figure 5 are samples of several cross sections that we examined). The lengths of the major and minor axis are about 4.0 and 1.5 km, respectively. The seismicity does not define a single plane, suggesting that the swarm earthquakes are not aftershocks on a single fault.

Fault Mechanisms. Using the computer program FPFIT [Reasenber and Oppenheimer, 1985] we obtained fault mechanisms for 30 earthquakes in the San Acacia swarm which had seven or more first motion readings and latitude, longitude and depth errors ≤ 0.5 km. Twenty mechanisms are normal faults and strike roughly north (Figure 6). First motions for 10 of the events did not fit any standard double-couple source model (Figure 7). The ten events appear to have reduced dilatational quadrants, which may indicate rapid tensile opening as well as shear motions during faulting. We speculate that movement of groundwater in a geothermal reservoir could account for these observations, however further modeling is planned to help us better understand this earthquake swarm.

B. Summary and Conclusions

Important characteristics of the San Acacia swarm that have been revealed by this research are listed below:

1. The hypocenters occupy an ellipsoidal volume with major axis striking ENE and plunging 45° along strike.
2. The standard model with station corrections is an adequate velocity model for locations and fault mechanism analysis of swarm events, since the shallowest earthquakes are below the Phanerozoic-Precambrian interface.
3. Focal mechanisms for 20 of 30 events analyzed are normal and strike north. The other 10 events have first motions which do not appear to fit any standard double-couple.

We have not completed any source modeling at the time of this report, however, it is possible that groundwater circulating in heated rocks could allow the disparate focal mechanisms. Regardless, some additional component to the standard double-couple without moment source mechanism is necessary to explain the reduction in the dilatational quadrants of one-third of the well-constrained focal mechanisms.

C. Equipment Obtained and Facilities Used

No new equipment has been obtained for this project. Data analysis was completed at the New Mexico Tech Geophysics Computer Facility.

D. Tangible Results

Papers in Preparation

Sanford, A., R. Balch, J. Lakings, H. Hartse, and L. House. A link between listric faulting and recent seismicity in the central Rio Grande rift of New Mexico, in preparation for submission to Bulletin of the Seismological Society of America.

Balch, R., A. Sanford, L. Jaksha, H. Hartse, and L. House. Non-Double Couple Focal Mechanisms Observed in the Central Rio Grande Rift of New Mexico, in preparation for submission to Bulletin of the Seismological Society of America.

Abstracts and Presentations

Sanford, A., R. Balch, J. Lakings, H. Hartse, and L. House, A link between listric faulting and recent seismicity in the central Rio Grande rift of New Mexico, abstract presented at Fall, 1993 AGU Meeting.

Balch, R., A. Sanford, L. Jaksha, H. Hartse, and L. House. Non-Double Couple Focal Mechanisms Observed in the Central Rio Grande Rift of New Mexico, abstract submitted to Fall, 1994 AGU Meeting.

E. Intangible Results

The administration of this project has encouraged collaboration between earth scientists at New Mexico Tech and LANL. Communication through e-mail has been routine. In addition, Hans Hartse has a computer account on the New Mexico Tech Geophysics Sun network. From his office in Los Alamos

he remotely logs on to computers at the New Mexico Tech Computer Facility, where he installs upgraded software and assists collaborators at New Mexico Tech.

III. Future of the Project

For FY95 we have been funded to expand our study to include a series of earthquake swarms located on the eastern margin of the central Rio Grande Rift. The initial part of the study will involve obtaining accurate hypocenters and focal mechanisms following techniques used in the previous two studies. The continuing goal of the project is to study and characterize microearthquake swarms in continental rifts and the relationship between rifting and swarm activity, whether direct or indirect.

IV. Funds Received

For FY94 we received \$25,000, with \$20,000 going to New Mexico Tech and \$5000 going to LANL (EES-4). Funds at New Mexico Tech support Robert Balch (a doctoral student) and an undergraduate research assistant and partially support the Geophysics Computer Facility at Socorro. Funds at LANL partially support Hans Hartse, a post-doc with EES-4. Our funding has been extended for FY95 at a level of \$25,000, with \$20,000 going to New Mexico Tech.

V. Los Alamos Facilities Used

Hans Hartse and Leigh House use the EES-4 Sun network for event processing and e-mail/ftp communications with researchers at New Mexico Tech.

VI. References

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- Larsen, S., R. Reilinger, and L. Brown (1986). Evidence of ongoing crustal deformation related to magmatic activity near Socorro, New Mexico, *J. Geophys. Res.*, **91**, 6283-6292.
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- Rinehart, E. J., A. R. Sanford, and R. M. Ward (1979)., Geographic extent and shape of an extensive magma body at mid-crustal depths in the Rio Grande rift near Socorro, New Mexico, in *Rio Grande Rift: Tectonics and Magmatism*, edited by R. E. Riecker, pp. 237-251, AGU, Washington, D. C.

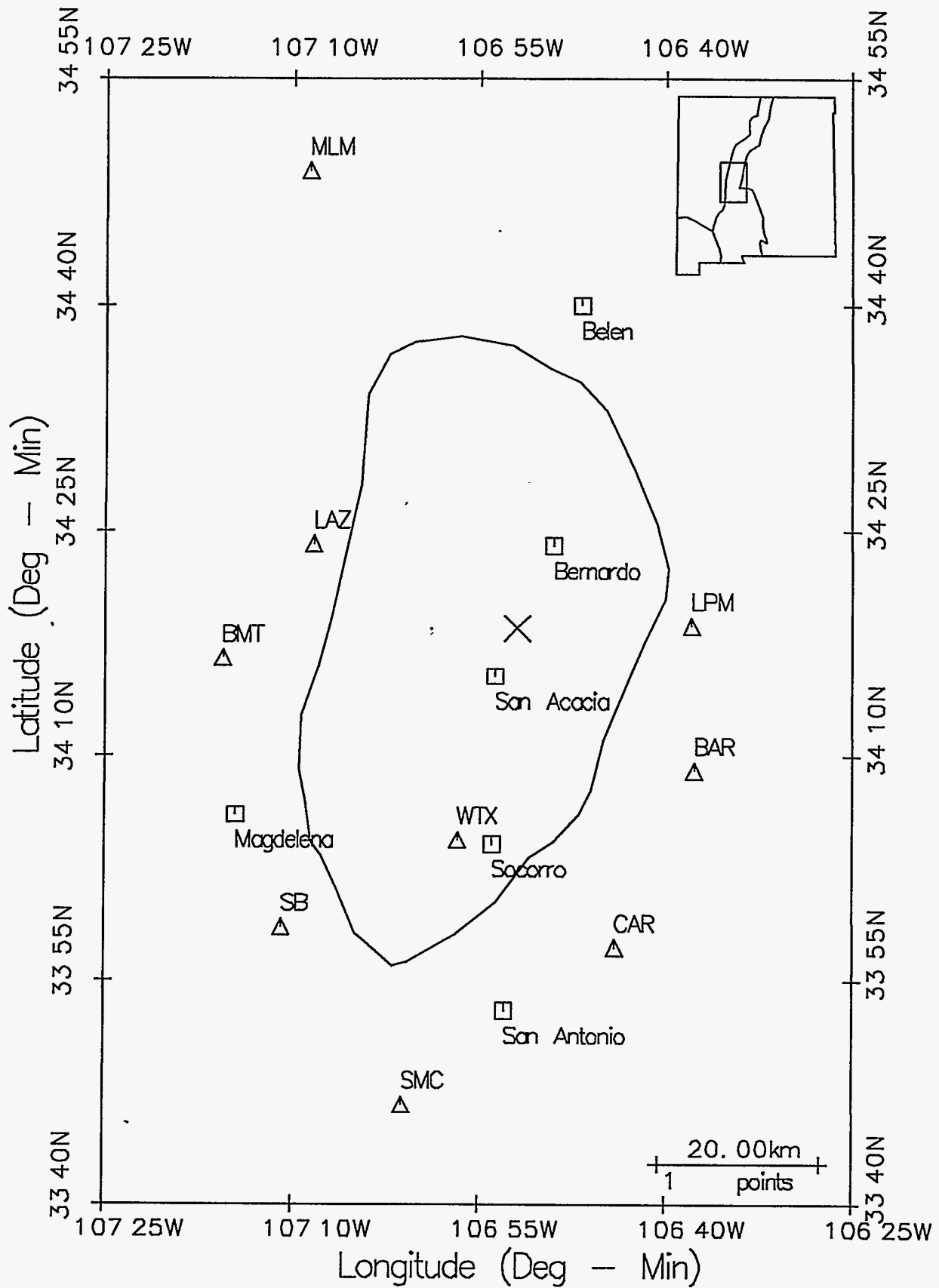


Figure 1. Seismic network operated near Socorro, NM by New Mexico Tech. Triangles indicate permanent station sites and the outline represents the Socorro Magma Body from Hartse and Sanford [1992]. The large X shows the San Acacia swarm epicenter.

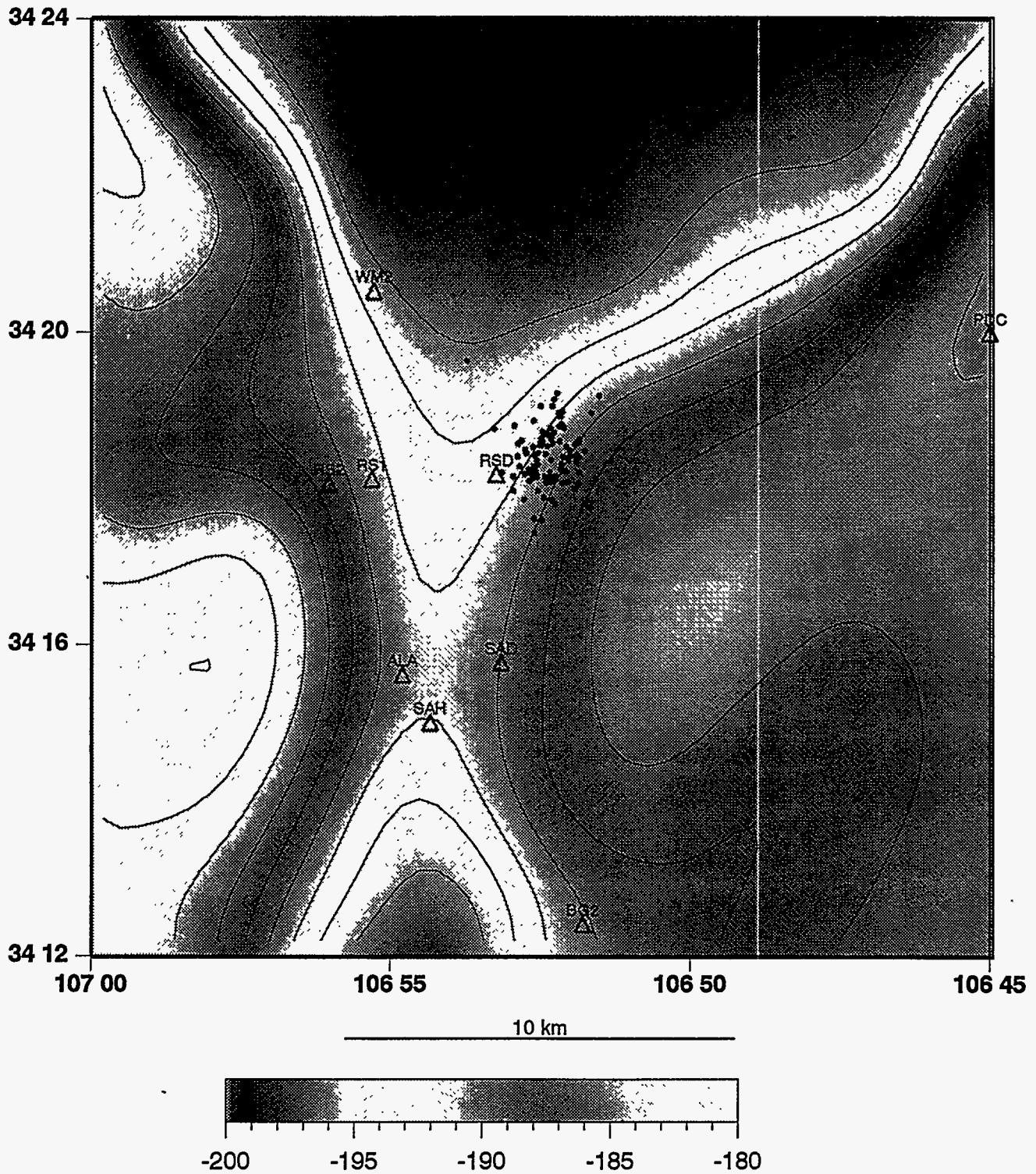


Figure 2. Bouguer anomaly contour map for the San Acacia, New Mexico area. Also shown are the locations of temporary station sites (triangles) and the 102 best epicenters (black circles). Color roughly correlates to sediment thickness, white and green being thinnest and thickest sediments, respectively.

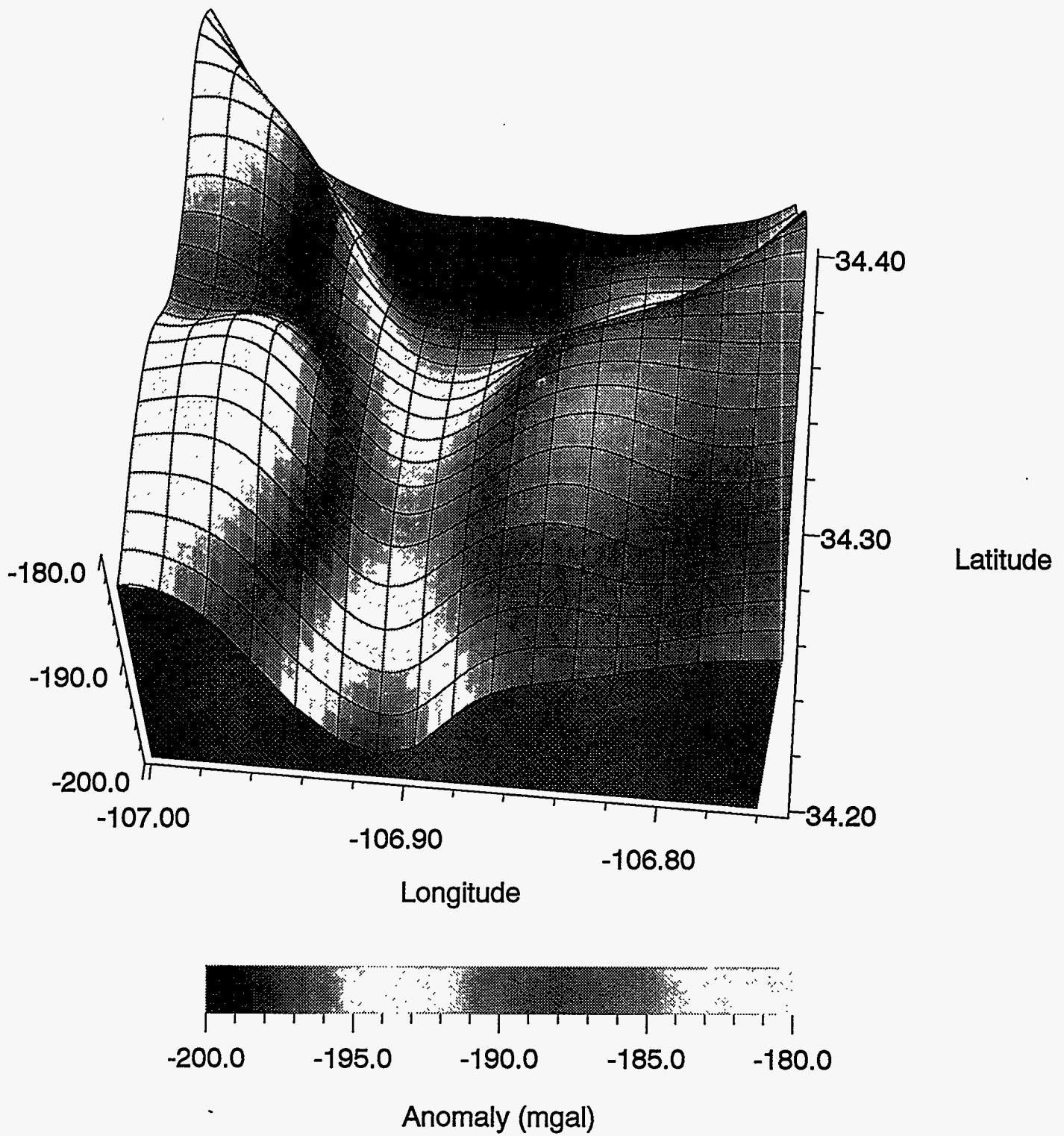


Figure 3. Three-dimensional representation of the contour plot in Figure 2, further illustrating the saddle between two RGR basins. Green correlates with the thickest sediment layers and white the thinnest.

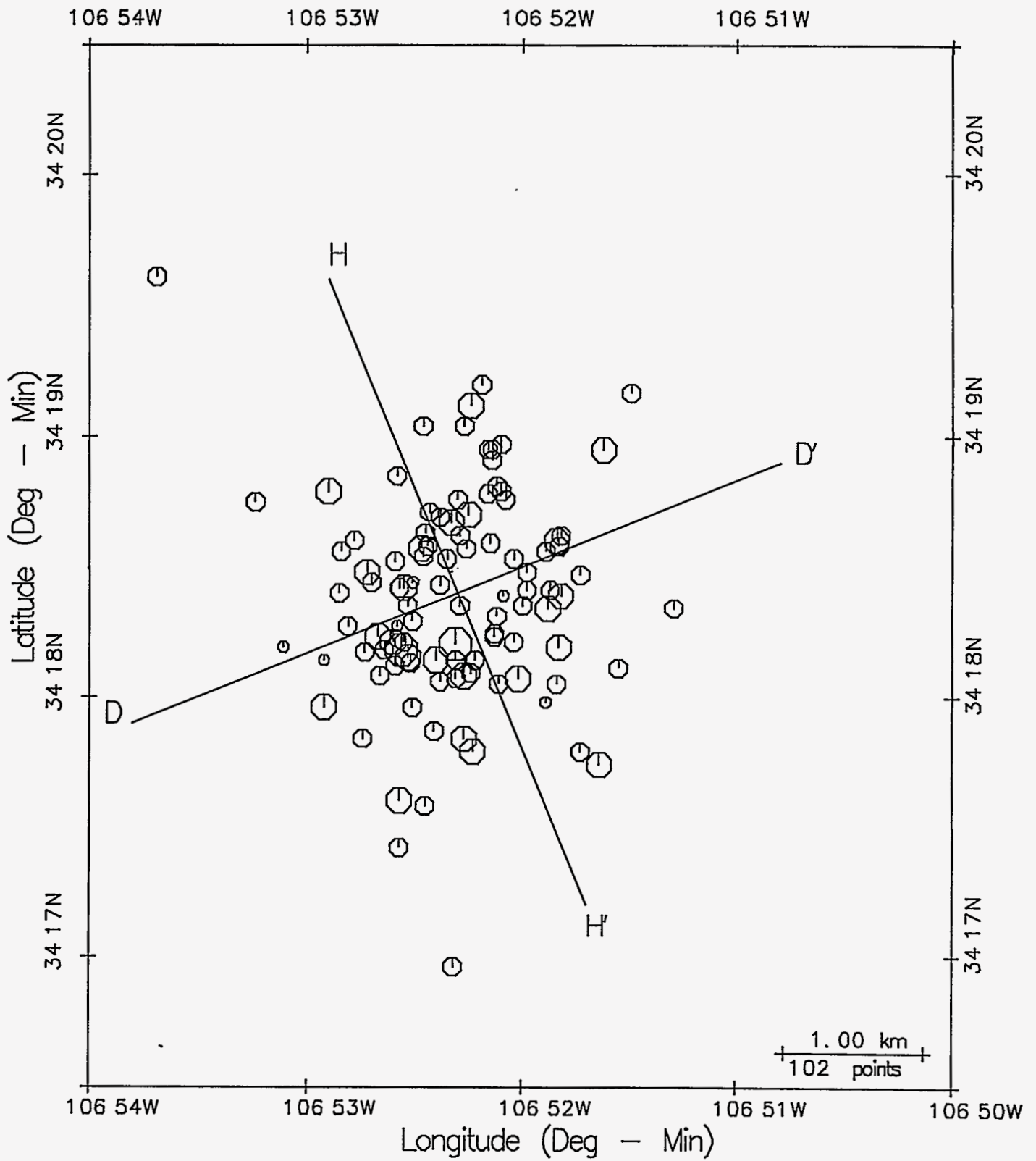


Figure 4. Epicenters for earthquakes in the San Acacia swarm. The map is restricted to the 102 events that have errors in latitude, longitude and depth ≤ 0.5 km.

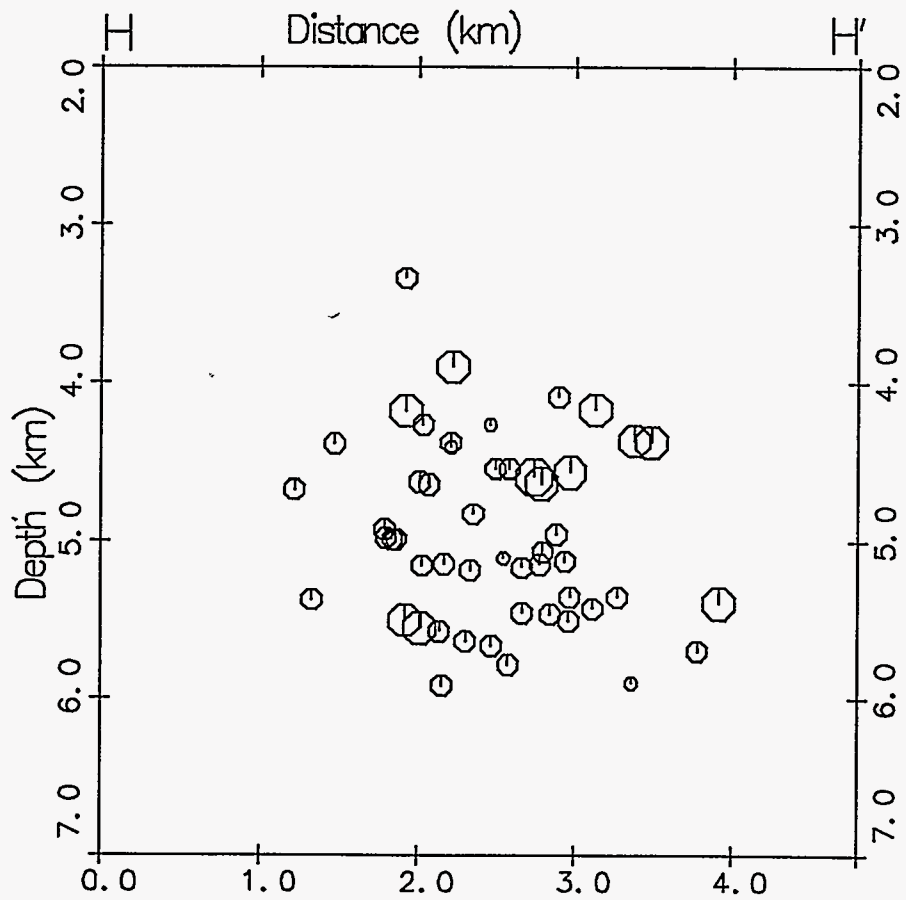
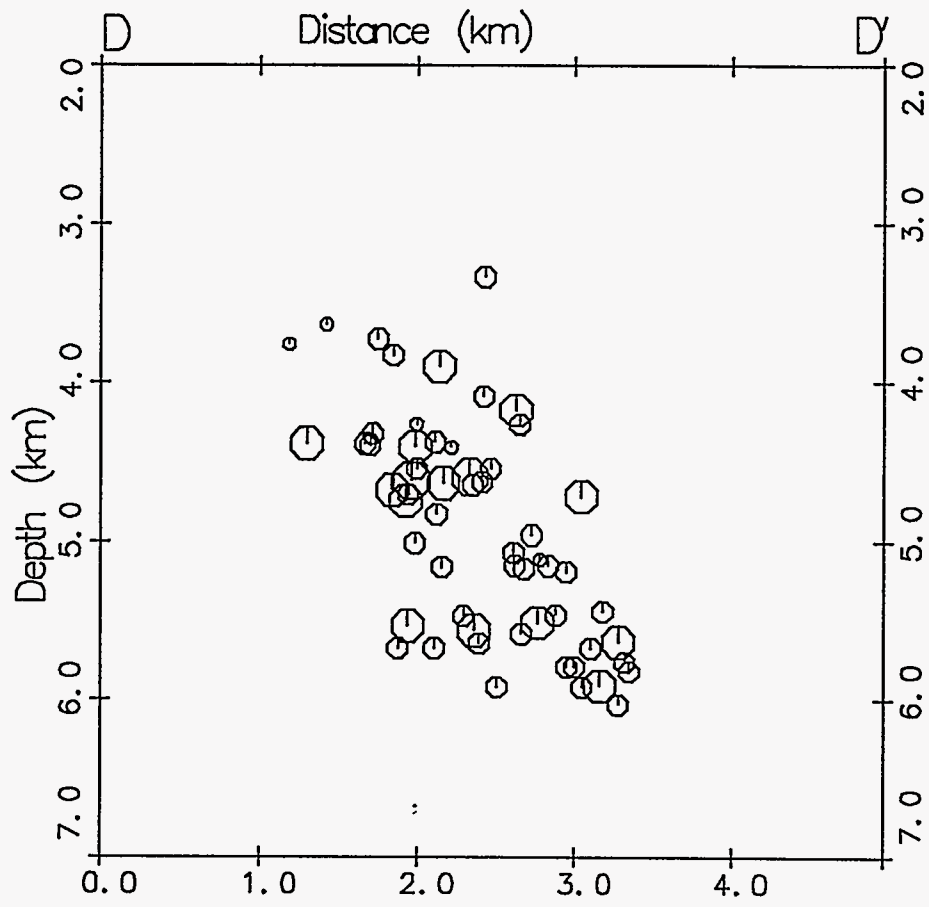


Figure 5. Cross-sections DD' and HH' from Figure 4.

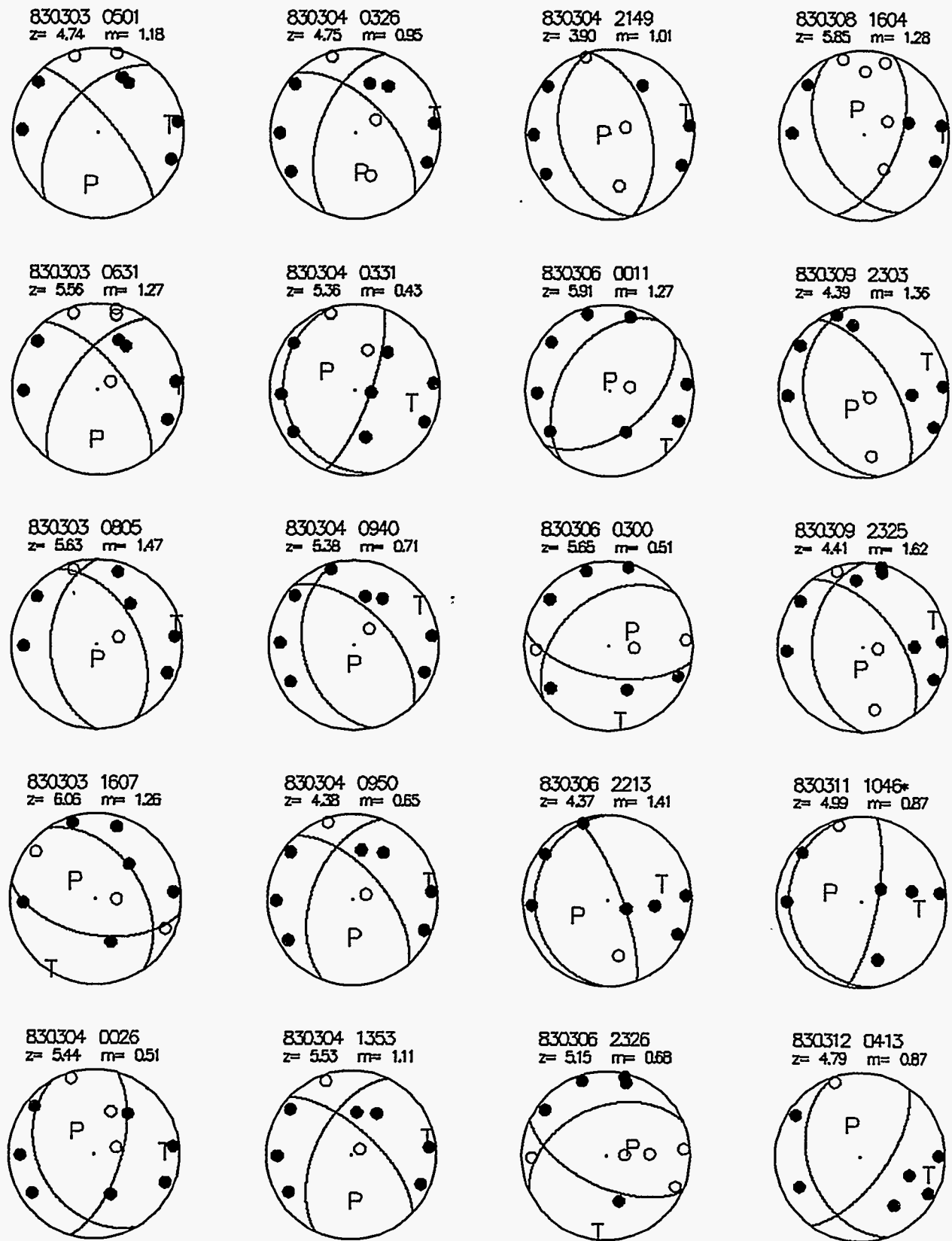


Figure 6. Fault mechanisms constrained to fit double-couple sources. Most can be constrained as normal faults striking roughly north.

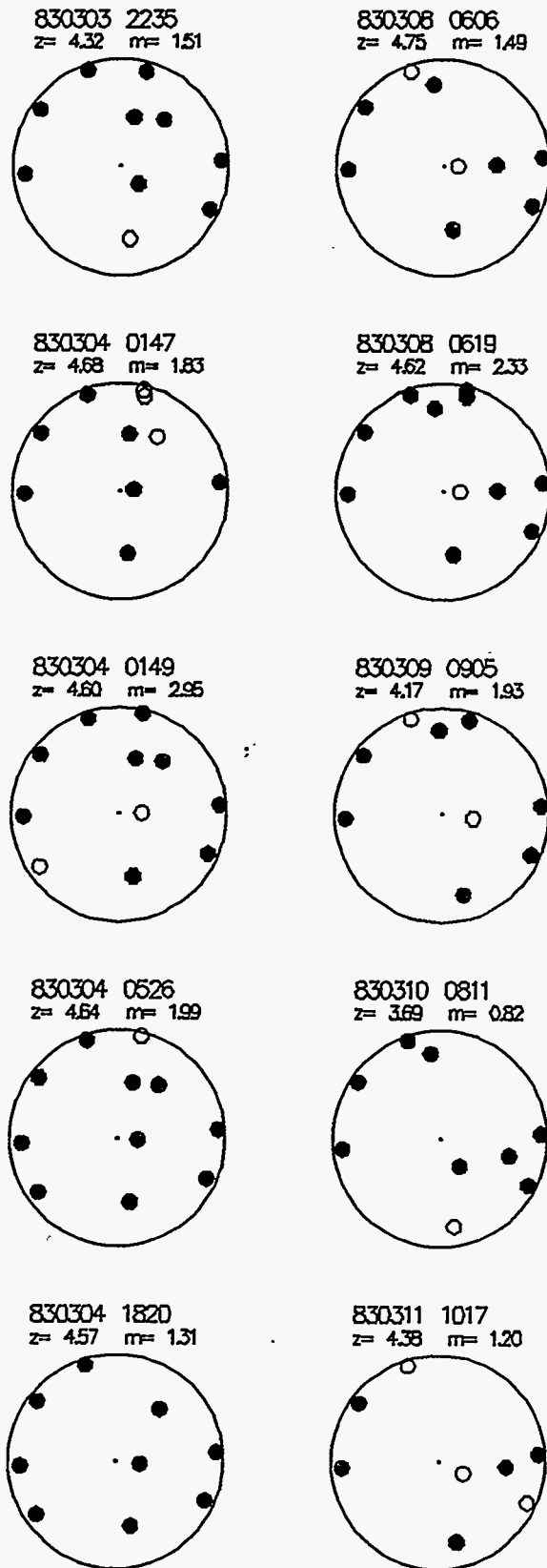


Figure 7. Fault mechanisms which do not fit standard double-couple sources.
Note the reduced dilatational quadrants.

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