

1 **Vertical stratification of testate amoebae in the Elatia Mires, northern Greece:**
2 **Palaeoecological evidence for a wetland response to recent climatic change, or**
3 **autogenic processes?**

4

5 Richard J. Payne^{1,2} and Jacqueline M. Pates³

6

7 ¹ The Fitch Laboratory, British School at Athens, Odhós Soudhías 52, GR-10676 Athens,
8 Greece.

9 ² Geography, School of Environment and Development, The University of Manchester,
10 Oxford Road, Manchester, M13 9PL, UK.*

11 ³Department of Environmental Science, Lancaster Environment Centre, Lancaster
12 University, Lancaster, LA1 4YQ, UK.

13 * To whom correspondence should be addressed. E-mail: r.j.payne@manchester.ac.uk

14

15 ABSTRACT

16

17 The Elatia Mires of northern Greece are unique ecosystems of high conservation value.
18 The mires are climatically marginal and may be sensitive to changing hydroclimate,
19 while northern Greece has experienced a significant increase in aridity since the late 20th
20 century. To investigate the impact of recent climatic change on the hydrology of the
21 mires, the palaeoecological record was investigated from three near-surface monoliths
22 extracted from two sites. Testate amoebae were analysed as sensitive indicators of
23 hydrology. Results were interpreted using transfer function models to provide

1 quantitative reconstructions of changing water table depth and pH. AMS radiocarbon
2 dates and ^{210}Pb suggest the peats were deposited within the last c.50 years, but do not
3 allow a secure chronology to be established. Results from all three profiles show a
4 distinct shift towards a more xerophilic community particularly noted by increases in
5 *Euglypha* species. Transfer function results infer a distinct lowering of water tables in this
6 period. A hydrological response to recent climate change is a tenable hypothesis to
7 explain this change; however other possible explanations include selective test decay,
8 vertical zonation of living amoebae, ombrotrophication and local hydrological change. It
9 is suggested that a peatland response to climatic change is the most probable hypothesis,
10 showing the sensitivity of marginal peatlands to recent climatic change.

11

12 KEYWORDS: Mires, Peatlands, Climate Change, Testate Amoebae, Palaeohydrology

13

1 INTRODUCTION

2 Climate change may lead to changes in peatland carbon sequestration (Belyea and
3 Malmer, 2004; Lavoie et al., 2005), gas flux (Chapman and Thurlow, 1998; Regina et al.,
4 1999; Keller et al., 2004), plant communities (Weltzin et al., 2003; Robroek et al., 2007)
5 and permafrost melting (Camill, 2005; Turetsky et al., 2007). In many regions of the
6 world impacts on peatland hydrology have been noted (e.g. Klein et al., 2005). The
7 peatlands likely to show the greatest impacts are the most marginal sites in regions which
8 experience the greatest climate change. Such a case may be the Mediterranean region,
9 where peatlands are rare but are often climatically marginal and are likely to be sensitive
10 to comparatively minor climatic changes. Mediterranean peatlands are relatively
11 unimportant in terms of their carbon reserve but are extremely important in terms of
12 regional biodiversity.

13 Meteorological records for northern Greece show a general trend of reduced
14 precipitation over the last c.50 years. Climatic impacts on the hydrology of northern
15 Greece have already been noted and are modelled to increase in severity over coming
16 decades with consequent impacts for human populations (Mimikou, 1993; Mimikou et
17 al., 1999; 2000; Bürger, 2002; Baltas and Mimikou, 2005). These changes may have
18 affected peatlands in the region. Most peatlands in Greece have been heavily impacted by
19 human activity primarily through drainage for agriculture in the 20th Century, and often
20 have low conservation value (Bouzinou et al., 1994; 1997; Christanis, 1996). An
21 exception to this is the Elatia Mires of northern Macedonia, where remoteness and Cold
22 War access restrictions have prevented major impacts. These mires are the most
23 oligotrophic peatlands in Greece and are probably the only location in the southern

1 Balkans where many wetland endemics typical of more nutrient-poor conditions may be
2 found. These are the only peatlands in Greece where *Sphagnum* is a permanent presence.
3 Due to their unique nature the Elatia Mires have been a focus of recent conservation
4 attention. This study uses a palaeoecological approach based on testate amoebae analysis
5 to test the hypothesis that the hydrology of the Elatia mires is responding to recent
6 climatic changes.

7

8 SITES and METHODS

9

10 The Elatia Mires lie in the Elatia Forest, approximately 70 km north of the city of
11 Drama and 5 km south of the Bulgarian frontier at around 1500m asl. (41°29'N, 24°19'E;
12 Fig. 1). Four small peatlands are situated in clearings within a natural coniferous forest
13 dominated by *Picea abies*, the only such forest in Greece (Papazisimou et al., 2002). A
14 comparatively cool and moist climate combined with impermeable granite bedrocks have
15 allowed peat formation. Peat deposits are shallow, probably not exceeding 1 m of
16 continuous peat, although deeper, buried, peat deposits are also found in one site. The
17 mires have been termed the sites 'transitional' due to their mix of features typical of fens
18 and bogs (Papazisimou et al 2002). Samples were extracted from the two most
19 oligotrophic sites: Dexameni mire (site DE; mean pH 6.5) and Krya Vrissi mire (site KB;
20 mean pH 6.4). Vegetation of the mires includes *Juncus effusus*, *Carex* spp., *Eriophorum*
21 *latifolium*, *Myosotis* spp., *Geum* spp., *Ranunculus* spp., *Mentha spicata*, *Plagiomnium*
22 *elatum*, *Sphagnum flexuosum*, *Aulacomnium palustris* and *Climacium dendroides*
23 (Papazisimou et al 2002).

1 Peat monoliths between 16 and 26 cm in length were extracted by cutting down
2 from the peat surface. Two monoliths were removed from central areas of the Krya Vrissi
3 mire and one from Dexameni; denoted KB1, KB2 and DE respectively. Peat stratigraphy
4 was noted and humification recorded on the Von Post scale (Von Post, 1924).

5 Testate amoebae analysis was used to reconstruct changes in hydrology through
6 the length of the monoliths. Testate amoebae are shell-forming unicellular
7 microorganisms that are abundant in peatlands and sensitive to peatland hydrology. By
8 analysing the changing community composition down the length of a peat core and
9 interpreting the results with a transfer function model it is possible to reconstruct how
10 mire wetness has varied over time (Charman, 2001; Mitchell et al., 2008). Such
11 reconstructions have been validated by comparison with instrumental data and
12 independent proxy-climatic records (Charman and Hendon, 2000; Charman et al., 2004;
13 Schoning et al., 2005).

14 Peat sub-samples for testate amoebae analysis were boiled in deionised water,
15 filtered at 250 μm and then back-filtered at 15 μm with the 15 $>$ 250 μm fraction retained
16 (Hendon and Charman, 1997). Slides were made up with glycerol and amoebae identified
17 following the taxonomic scheme described in Payne and Mitchell (2007); a count of 150
18 tests was aimed for. Amoebae diagrams were constructed using C² ver. 1.4 (Juggins,
19 2003) and zoned using optimal sum of squares partitioning (Birks and Gordon, 1985) in
20 ZONE ver. 1.2 (Juggins, 1992). Quantitative environmental reconstruction was carried
21 out using the transfer function model developed by Payne and Mitchell (2007). Depth to
22 water table (DWT) was reconstructed using a maximum likelihood model (RMSEP_{boot}
23 1.9cm) and pH using a weighted average model (RMSEP_{boot} 0.4). The reconstructed

1 values are termed testate amoebae inferred depth to water table (TI-DWT) and testate
2 amoebae inferred pH (TI-pH). Bootstrapped error estimation with 1000 cycles was used
3 to provide sample-specific error estimates.

4

5 Chronology

6 Establishing reliable chronologies for recent peat deposits has been a persistent
7 challenge for peatland palaeoecologists (Turetsky et al., 2004) and is particularly
8 problematic for these sites as they are non-ombrotrophic and include unusual plant
9 communities. A search for cryptotephra (following the method of Pilcher and Hall,
10 1992) failed to identify any shards and a search for spheroidal carbonaceous particles
11 (following the method of Rose et al, 1995) failed to find adequate concentrations. Two
12 radiometric methods, ^{210}Pb and ^{14}C analysis were applied to the peat profiles. ^{210}Pb was
13 analysed by assuming equilibrium with its grand-daughter ^{210}Po . Peat samples were
14 dissolved in strong acids with a ^{209}Po yield tracer, plated onto copper disks and activity
15 measured by α -spectrometry. Four samples from towards the base of the monoliths were
16 AMS radiocarbon dated (Goodsite et al, 2001; Goslar et al, 2005). Bulk samples were
17 used due to the absence of *Sphagnum* (Nilsson et al. 2001). Samples were carefully
18 prepared to minimise risk of external contamination and rootlets were picked out.

19

20 RESULTS

21 Testate amoebae

22 Testate amoebae were found through the length of the monoliths but apparent
23 concentrations were low (counting time approximately 6-8 hours per sample). The three

1 testate amoebae diagrams show similar trends (Fig. 2). At the base of the profiles (zones
2 DE-1, KB1-1 and the lower portion of KB2-1) the community composition is noted by
3 abundant *Diffflugia* spp., particularly *Diffflugia pulex* type. Above this section there is a
4 decline in these taxa leading to a more diverse community noted by *Centropyxis*
5 *aerophila* (zones DE-2, KB1-2 and the upper portion of KB2-1). The next significant
6 change is a marked increase in *Euglypha rotunda*, and in KB1, *Euglypha ciliata* type. The
7 uppermost samples are different from those directly below, recognized as a separate zone
8 in KB2 (zone KB2-3) and DE (zone DE-3).

9 Overlap between the palaeoecological data and the modern training set is very
10 good. Over 99% of all amoebae counted are included in the training set; the total for
11 individual samples does not fall below 97%. Water table reconstructions show a similar
12 pattern between sites (Fig. 3). From the base of the sequences to c.6 cm depth there is low
13 amplitude variability with no obvious similarity between profiles. At 6 cm there is a rapid
14 increase in values that significantly exceeds bootstrapped error estimates, representing a
15 significant lowering of water table in all sites. At the top of the sequence there is a slight
16 decline in TI-DWT values and then a resumed increase in sites KB1 and DE, although
17 there is a continued decline in site KB2. pH reconstructions show gradual lowering of pH
18 values, increasing in rate above c.6 cm. The changes in pH are less pronounced than
19 those in DWT and (particularly in KB2) only marginally exceed the error estimates.

20 The main amoebae community change is a shift in the most abundant taxon to
21 *Euglypha rotunda* from *Centropyxis aerophila* type and *Diffflugia pulex* type. *C.*
22 *aerophila* is generally regarded as typical of moderately wet conditions, while *E. rotunda*
23 is probably most typical of intermediate conditions (Charman et al., 2000). The indicator

1 value of *D. pulex* type is uncertain. The majority of transfer functions to encounter both
2 taxa show *E. rotunda* to have a higher DWT optimum than *C. aerophila* type (or
3 equivalent *C. cassis* type) (Payne et al., 2006; 2007; in press; Woodland et al., 1998;
4 Charman and Warner, 1997; Charman, 1997; Warner and Charman, 1994; Lamentowicz
5 and Mitchell, 2005). The transfer function results therefore agree with the known
6 autecology of the taxa in interpreting this change as a shift to drier conditions.

7

8 Chronology

9 The ^{210}Pb results show no decline in activity with depth and a sequence of peaks
10 and troughs (Fig. 4). It was not possible to determine the “unsupported” component of the
11 ^{210}Pb ; insufficient material was available for direct ^{226}Ra analysis and the base of the
12 unsupported layer was not reached. The lack of a monotonic decrease in activity with
13 depth undermines the attempt to use the method for dating. There are two possible causes
14 of these results; very rapid peat accumulation, as suggested by the radiocarbon dates, and
15 movement of lead within the peat profile given the minerotrophic nature of the sites
16 (MacKenzie et al., 1998).

17 Radiocarbon dates were all returned as post-bomb and calibrated using CaliBomb
18 (<http://calib.qub.ac.uk/CALIBomb/frameset.html>). All dates give multimodal probability
19 distributions ranging from the mid-1950s to post-1995, indicating rapid peat
20 accumulation (Table 1). As the dates were based on bulk samples it is possible that they
21 have been contaminated by modern carbon, perhaps through penetration by sedge roots.
22 However, the samples were prepared carefully to avoid contamination and obvious roots
23 were removed. Systematic differences between dates on bulk samples and selected

1 macrofossils have not been proven (Blaauw et al. 2004). The dates are internally
2 consistent in showing the peats to be late 20th century in age and are not contradicted by a
3 date of 100 ± 40 ¹⁴C yrs BP at 31-36 cm from a neighbouring site (Papazisimou et al.,
4 2002). For two samples from the KB2 monolith, the deeper sample (GdA-1178: 21-
5 24cm) has a highest probability peak more recent than the upper sample (GdA-1016: 13-
6 15cm). However this date also has a subsidiary probability peak at 1957-1958 (10.3%).
7 As there is no stratigraphic reason to suspect a reversal it seems more likely that this
8 older peak is the correct one.

9 Neither the ²¹⁰Pb results nor the radiocarbon dates allow us to establish a secure
10 chronology for these profiles. However both sets of results can be taken to suggest that
11 the sediments are very recent. Most probably these peats have accumulated within the last
12 few decades but it is not possible to be more precise. Nevertheless, there is no reason to
13 suspect disturbance of the stratigraphy and it is still probable that these profiles do
14 preserve a continuous record of testate amoebae changes and may reveal recent
15 hydrological changes in the mires.

16

17 DISCUSSION

18

19 The testate amoebae results show a single large change in amoebae community at
20 around 6 cm depth which can be inferred as a shift to a drier mire surface. This is entirely
21 consistent with the hypothesis of a lowered water table due to the general reduction in
22 precipitation in northern Greece. However simple attribution of the changes to a climatic

1 impact is not possible due to the presence of other factors which could also be responsible
2 for these changes.

3

4 1. Non-climatic hydrological change

5 It seems probable that the Elatia Mires have undergone hydrological change;
6 however as the sites are not ombrotrophic it is possible that this is unrelated to climate. It
7 is possible that processes such as forestry, tectonic or geomorphological change in the
8 wider area could have lead to a change in water input into the mires. Although there is no
9 particular evidence to suggest that this may be the case the hypothesis cannot be
10 discounted on the basis of the data presented here.

11

12 2. Vertical zonation of living amoebae

13 Interpretation of the uppermost testate amoebae assemblages may be complicated
14 if the amoebae are still alive below the surface. Testate amoebae have been noted to
15 exhibit vertical zonation forced by gradients in light, moisture and mineral material for
16 test construction (Heal, 1962; Booth 2002; Mitchell and Gilbert 2004; Mazei et al. 2007).
17 It is not clear that vertical zonation could explain the species changes observed. Taxa
18 with xenosome tests such as *Diffflugia* spp. and *C. aerophila* are commonly observed in
19 lower horizons, probably due to availability of material for test construction. However the
20 typical position of *Euglypha* species varies between studies (Chacharonis 1956; Booth
21 2002; Mitchell and Gilbert 2004; Mazei and Bobnova 2007). The only study in
22 transitional mires (Mazei and Bobnova 2007) found vertical zonation to be much weaker
23 than in bogs. In these sites alive or encysted amoebae were only noted in the top 2-3 cm.

1 It is probable that vertical zonation in this region could explain the unusual communities
2 in the uppermost samples, but unlikely that vertical zonation is the cause of the major
3 change at 6cm.

4

5 3. Test preservation

6 The taxa which are primarily responsible for the increase in TI-DWT in the upper
7 portions of the sequence (*E.rotunda*, *E.ciliata*, *C.dubium*) all have tests constructed of
8 idiosomes. Such tests may be particularly prone to decomposition in the fossil record
9 (Lousier and Parkinson, 1981; Swindles and Roe, 2007; Payne 2007). Selective test loss
10 could have led to inaccurate palaeoenvironmental reconstruction in these sites (Mitchell
11 et al. 2008). A number of strands of evidence suggest this is unlikely to be the major
12 cause of the changes: 1) during microscopy no apparent reduction in overall test
13 concentrations was noted with depth (although concentrations were not enumerated), this
14 might be expected given the high proportion of idiosome tests; 2) no increase in degraded
15 tests was noted with depth; 3) the taxa concerned decline sharply at around 6 cm,
16 however they continue through the rest of the profiles in lower concentrations; 4) other
17 closely related taxa reach their highest concentrations lower in the profiles (e.g.
18 *E.tuberculata* in KB1); 5) the timescale under consideration is very short so differential
19 preservation is perhaps less likely than in longer-term studies; 6) when Euglypha is
20 removed from the reconstructions an increase in TI-DWT at 6cm remains (Fig. 3),
21 although this is less marked and there are other changes such as a large drop in TI-DWT
22 at the top of KB1.

23

1 4. Autogenic mire development processes

2 The *Elatia* mires have been termed ‘transitional’ between fens and bogs; it is
3 possible that the changes in the palaeoecological record are due to ombrotrophication.
4 Ombrotrophication can be expected to lead to major amoebae community changes as fens
5 have distinctly different testate amoebae communities from bogs and nutrient status is an
6 important secondary gradient (Opravilova and Hajek, 2006).

7 Conventionally it has been assumed that an autogenic mechanism drives
8 ombrotrophication; peat gradually accumulates above the water table until it is no longer
9 reliant on groundwater and becomes acidified by leaching and the establishment of
10 *Sphagnum* (Hughes, 2000). This would sit well with the reconstructed increase in TI-
11 DWT and slight decrease in TI-pH. However, it seems unlikely that this autogenic model
12 could explain the suddenness of the amoebae change, the 2cm+ offset between
13 stratigraphic and amoebae community changes and the apparent synchronicity in change
14 between two sites. More recent studies have suggested allogenic forcing of
15 ombrotrophication with a lowered water table leading to the peat surface being separated
16 from groundwater (Hughes, 2000; Hughes and Dumayne-Peaty, 2002; Hughes and
17 Barber, 2003; Hughes et al., 2000). It is therefore also possible that ombrotrophication is
18 occurring, but is driven by a real allogenic change.

19

20 The results presented here are consistent with the hypothesis of a climate change-
21 induced hydrological change; however other explanations cannot be excluded. It is
22 possible that multiple processes have lead to the observed patterns. If the changes are due
23 to climate then this study provides the first evidence for the impacts of recent climate

1 change on Mediterranean peatlands. Climate change is likely to be a key challenge to
2 regional peatland conservation.

3

4 **Acknowledgements**

5

6 This study was primarily funded by a BSA MacMillan-Rodewald fellowship to RJP.
7 Radiocarbon dates were funded by grants from the Richard Bradford McConnell Fund
8 and a Gladstone Memorial Prize. Thanks to Kimon Christanis and Stavros Kalaitzidis
9 (University of Patras) for discussion of the Elatia Mires and to Edward Mitchell
10 (Lausanne) for discussion of testate amoebae ecology and taxonomy. The map was
11 prepared by Edward Mitchell. This work was carried out by permission of the Greek
12 Institute of Geology and Mineral Exploration (IGME) and the Forest Service of the
13 Prefecture of Drama.

14

15 **References**

16

17 Baltas EA, Mimikou MA (2005) Climate change impacts on the water supply of
18 Thessaloniki. *Water Resources Development* 21:341-353

19

20 Belyea L, Malmer N (2004) Carbon sequestration in peatland: patterns and mechanisms
21 of response to climate change. *Glob Change Biol* 10:1043-1052

22

- 1 Birks HJB, Gordon AD (1985) Numerical methods in Quaternary pollen analysis.
2 Academic Press, London
3
- 4 Blaauw M, van der Plicht J, van Geel B (2004) Radiocarbon dating of bulk peat samples
5 from raised bogs: non-existence of a previously reported ‘reservoir effect’? Quaternary
6 Sci Rev 23:1537-1542
7
- 8 Booth RK (2002) Testate amoebae as paleoindicators of surface-moisture changes on
9 Michigan peatlands: modern ecology and hydrological calibration. J Paleolimnol
10 28:329:348
11
- 12 Bouzinos A, Broussoulis J, Christanis K (1994) Conservation and management of Greek
13 fens: A “model” to avoid. In: Proceedings of the International Symposium: Conservation
14 and management of fens, Falenty Agricultural University, Warsaw
15
- 16 Bouzinos A, Christanis K, Kotis T (1997) The Chimaditida fen (W. Macedonia, Greece):
17 a peat deposit lost. Int Peat J 7:3-10
18
- 19 Bürger G (2002) Selected precipitation scenarios across Europe. J Hydrol 262: 99-110
20

1 Chapman S, Thurlow M (1998) Peat respiration at low temperatures. *Soil Biol Biochem*
2 30:1013-1021
3

4 Camill P (2005) Permafrost Thaw Accelerates in Boreal Peatlands During Late-20th
5 Century Climate Warming. *Climatic Change* 68:135-152
6

7 Charman D (1997) Modelling hydrological relationships of testate amoebae (Protozoa:
8 Rhizopoda) on New Zealand peatlands. *J Roy Soc New Zeal* 27: 465-83
9

10 Charman D (2001) Biostratigraphic and palaeoenvironmental applications of testate
11 amoebae. *Quaternary Sci Rev* 20: 1753-1764
12

13 Charman D, Warner B (1997) The ecology of testate amoebae (Protozoa: Rhizopoda) and
14 microenvironmental parameters in Newfoundland, Canada: modeling hydrological
15 relationships for palaeoenvironmental reconstruction. *Ecoscience* 4:555-562
16

17 Charman D, Hendon D (2000) Long-term changes in soil water tables over the past 4500
18 years: relationships with climate and North Atlantic atmospheric circulation and sea
19 surface temperatures. *Climatic Change* 47:45-59
20

21 Charman D, Hendon D, Woodland W (2000) The identification of testate amoebae
22 (Protozoa: Rhizopoda) from British oligotrophic peats. *Quaternary Research Association*
23 *Technical Guide Series*, Cambridge

1
2 Charman DJ, Brown AD, Hendon D, Kimmel A, Karofeld E (2004) Testing the
3 relationship between Holocene peatland palaeoclimate reconstructions and instrumental
4 data. *Quaternary Sci Rev* 23:137-143
5
6 Chacharonis P (1956) Observations on the ecology of protozoa associated
7 with *Sphagnum*. *J Protozool* 3:11
8
9 Christanis K (1996) The peat resources in Greece. In: Lappalainen E (ed) *Global Peat*
10 *Resources*. International Peat Society, Jyskä, pp87-90
11
12 Goodsite M, Rom W, Heinemeier J, Lange T, Ooi S, Appleby P, Shotyk W, van der
13 Knaap W, Lohse C, Hansen T (2001) High resolution AMS ¹⁴C dating of post-bomb peat
14 archives of atmospheric pollutants. *Radiocarbon* 43:495-515
15
16 Goslar T, van der Knaap WO, Hicks S, Andrič M, Czernik J, Goslar E, Räsänen S,
17 Hyötylä H (2005) Radiocarbon Dating of Modern Peat Profiles: Pre- and Post-Bomb ¹⁴C
18 Variations in the Construction Of Age-Depth Models. [Radiocarbon](#) 47: 115-134
19
20 Heal O (1962) The abundance and micro-distribution of testate amoebae
21 (Rhizopoda: Testacea) in *Sphagnum*. *Oikos* 13:35-47
22

1 Hendon D, Charman D (1997) The preparation of testate amoebae (Protozoa: Rhizopoda)
2 samples from peat. *Holocene* 7:199-205
3
4 Hughes PDM (2000) A reappraisal of the mechanisms leading to ombrotrophy in British
5 raised mires. *Ecol Lett* 3:7–9.
6
7 Hughes PDM, Dumayne-Peaty L (2002) Testing theories of mire development using
8 multiple successions at Crymlyn Bog, West Glamorgan, South Wales, UK. *J Ecol* 90:
9 456-471
10
11 Hughes PDM, Barber KE (2003) Mire development across the fen - bog transition on the
12 Teifi floodplain at Tregaron Bog, Ceredigion, Wales and a comparison with thirteen other
13 raised bogs. *J Ecol* 91:253-264
14
15 Hughes PDM, Mauquoy D, Barber KE, Langdon PE (2000) Mire-development pathways
16 and palaeoclimatic records from a full Holocene peat archive at Walton Moss, Cumbria,
17 England. *Holocene* 10:465–479
18
19 Juggins S (1992) The ZONE program, version 1.2 (unpublished program). University of
20 Newcastle, Newcastle upon Tyne
21
22 Juggins S (2003) C2 user guide. Software for ecological and palaeoecological data
23 analysis and visualisation. University of Newcastle, Newcastle Upon Tyne

1
2 Keller J, White J, Bridgham S, Pastor J (2004) Climate change effects on carbon and
3 nitrogen mineralization in peatlands through changes in soil quality. *Glob Change Biol*
4 10:1053-1064
5
6 Klein E, Berg EE, Dial R (2005) Wetland drying and succession across the Kenai
7 Peninsula Lowlands, south-central Alaska. *Can J Forest Res* 35: 1931-41
8
9 Lamentowicz M, Mitchell EAD (2005) The ecology of testate amoebae (Protists) in
10 *Sphagnum* in north-west Poland in relation to peatland ecology. *Microbial Ecol* 50: 48-63
11
12 Lavoie M, Paré D, Bergeron Y (2005) Impact of global change and forest management
13 on carbon sequestration in northern forested peatlands. *Environ Rev* 13:199-240
14
15 Louiser J, Parkinson D (1981) The disappearance of the empty tests of litter and soil
16 testate amoebae (Testacea, Rhizopoda, Protozoa). *Arch Protistenkd* 124: 312-336
17
18 MacKenzie AB, Logan EM, Cook GT, Pulford ID (1998) Distributions, inventories and
19 isotopic composition of lead in ²¹⁰Pb-dated peat cores from contrasting biogeochemical
20 environments: Implications for lead mobility. *Sci Total Environ* 223: 25-35
21

1 Mazei YA, Bobnova OA (2007) Species composition and structure of testate amoebae
2 community in a *Sphagnum* Bog at the initial stage of its formation. Biology Bulletin 34:
3 619:628
4
5 Mazei YA, Tsyganov AN, Bubnova OA (2007) Structure of a community of testate
6 amoebae in a *Sphagnum* dominated bog in Upper Sura Flow (Middle Volga Territory).
7 Biology Bulletin 34: 382-394
8
9 Mimikou MA (1993) Extreme variations of the hydrological cycle in Greece. Variability
10 or change? Tech Chron, 13:67-81
11
12 Mimikou MA, Kanellopoulou SP, Baltas EA (1999) Human implication of changes in
13 the hydrological regime due to climate change in Northern Greece. Global Environ
14 Chang 9:139-156
15
16 Mimikou MA, Baltas E, Varanou E, Pantazis K (2000) Regional impacts of climate
17 change on water resources quantity and quality indicators. J Hydrol 234:95–109
18
19 Mitchell EAD, Gilbert D (2004) Vertical micro-distribution and response to nitrogen
20 deposition of testate amoebae in *Sphagnum*. J Eukaryot Microbiol 51:480-490
21

1 Mitchell EAD, Charman DJ, Warner BG (2008) Testate amoebae analysis in ecological
2 and paleoecological studies of wetlands: past, present and future. *Biodivers Conserv*
3 17:329–343
4
5 Mitchell E, Payne R, Lamentowicz M (2008) Potential implications of differential
6 preservation of testate amoebae shells for paleoenvironmental reconstruction in
7 peatlands, *J Paleolimnol* (in press) DOI 10.1007/s10933-007-9185-z
8
9 Nilsson M, Klarqvist M, Bohlin E, Possnert G (2001) Variation in ¹⁴C age of
10 macrofossils and different fractions of minute peat samples dated by AMS. *Holocene*
11 11:579:586
12
13 Opravilová V, Hájek M (2006) The variation of testacean assemblages (Rhizopoda) along
14 the complete base-richness gradient in fens: A case study from the western Carpathians.
15 *Acta Protozool* 45:191-204
16
17 Papazisimou S, Bouzinos A, Christanis K, Tzedakis PC, Kalaitzidis S (2002) The upland
18 Holocene transitional mires of Elatia forest, Northern Greece. *Wetlands* 22:355-365
19
20 Payne R, Kishaba K, Blackford J, Mitchell E (2006) The ecology of testate amoebae in
21 southcentral Alaskan peatlands: Building transfer function models for
22 palaeoenvironmental inference. *Holocene* 16:403-414
23

- 1 Payne R (2007) Laboratory experiments on testate amoebae preservation in peats:
2 implications for palaeoecology and future studies. *Acta Protozool* 46:325-332
3
- 4 Payne R, Mitchell E (2007) Ecology of testate amoebae from mires in the Central
5 Rhodope Mountains, Greece and development of a transfer function for
6 paleohydrological reconstruction. *Protist* 158:159-171
7
- 8 Payne R, Charman D, Eastwood W (2008) Testate amoebae as palaeoclimatic proxies in
9 Sürmene Ağaçbaşı Yaylasi peatland (Northeast Turkey). *Wetlands* (in press)
10
- 11 Pilcher J, Hall V (1992) Towards a tephrochronology for the Holocene of the north of
12 Ireland. *Holocene* 2:255-259
13
- 14 Regina K, Silvola J, Martikainen P (1999) Short-term effects of changing water table on
15 N₂O fluxes from peat monoliths from natural and drained boreal peatlands. *Glob Change*
16 *Biol* 5:183-189.
17
- 18 Robroek B, Limpens J, Breeuwer A, Crushell P, Schouten M (2007) Interspecific
19 competition between *Sphagnum* mosses at different water tables. *Funct Ecol* 21:805-812
20
- 21 Rose N, Harlock S, Appleby P, Battarbee R (1995) Dating of recent lake sediments in the
22 United Kingdom and Ireland using spheroidal carbonaceous particle (SCP) concentration
23 profiles. *Holocene* 5:328-335

1
2 Schoning K, Charman DJ, Wastegård S (2005) Reconstructed water tables from two
3 ombrotrophic mires in eastern central Sweden compared with instrumental
4 meteorological data. *Holocene* 15:111-118
5
6 Swindles GT, Roe HM (2007). Examining the dissolution characteristics of testate
7 amoebae (Protozoa: Rhizopoda) in low pH conditions: Implications for peatland
8 palaeoclimate studies. *Palaeogeogr Palaeocl* 252:486-496
9
10 Turetsky MR, Manning SW, Wieder RK (2004) Dating recent peat deposits. *Wetlands*
11 24:324-326
12
13 Turetsky M.R, Wieder RK, Vitt DH, Evans RJ, Scott KD (2007) The disappearance of
14 relict permafrost in boreal north America: Effects on peatland carbon storage and fluxes.
15 *Glob Change Biol* 13:1922-1934
16
17 Von Post L (1924) *Das genetische System der organogenen Bildungen Schwedens.*
18 *Comité International de Pedologie IV Commission, Helsinki*
19
20 Warner B, Charman D (1994) Holocene changes on a peatland interpreted from testate
21 amoebae (Protozoa) analysis. *Boreas* 23:270-280
22

1 Weltzin J, Bridgham S, Pastor J, Chen J, Harth C (2003) Potential effects of warming and
2 drying on peatland plant community composition. *Glob Change Biol* 9:141-151

3

4 Woodland W, Charman D, Simms P (1998) Quantitative estimates of water tables and
5 soil moisture in Holocene peatlands from testate amoebae. *Holocene* 8: 261-273

6

7

1

2 FIGURES and TABLE

3

4 Figure 1. Location of the Elatia Mires.

5

6 Figure 2. Testate amoebae profiles from three monoliths. Showing peat stratigraphy,
7 major testate amoebae species (% of total), testate amoebae total count, humification
8 expressed on the Von Post scale, and testate amoebae zones. Stratigraphic columns show
9 moss-dominated peat (solid wavy lines), peat of mixed composition with mosses most
10 abundant (interrupted wavy lines), peat of mixed composition with macrofossils most
11 abundant (interrupted vertical lines) and macrofossil-dominated peat (solid vertical lines).

12

13 Figure 3. Testate amoebae inferred depth to water table (TI-DWT) and testate amoebae
14 inferred pH (TI-pH) from the three peat monoliths with boot-strapped error estimates.
15 The transfer function model was based on one-off DWT and pH measurements so the
16 units of reconstruction are depth to water table (cm) and pH based on an October 2005
17 datum. Also showing TI-DWT and TI-pH reconstructions excluding *Euglypha* species,
18 which might be lost from the palaeoecological record.

19

20 Figure 4. Radiocarbon dates and ^{210}Pb profile for monolith KB2.

21

22

23

1 Table 1. Radiocarbon dates from the peat monoliths

2

Monolith	Depth (cm)	Date code	Percent modern carbon (pMc)	Calibrated age range (95% probability) (cal. years AD)
DE	15-16	GdA-1177	120.55 ± 0.37	1958-1961 (9.3%) 1985-1988 (90.5%)
KB1	21-24	GdA-1015	121.45 ± 0.36	1958-1961 (52.7%) 1984-1986 (42.3%)
KB2	13-15	GdA-1016	116.17 ± 0.33	1956-1958 (0.2%) 1989-1991 (94.8%)
KB2	21-24	GdA-1178	110.88 ± 0.53	1957-1958 (10.3%) 1995- (85.1%)

3