

## Sr and Nd isotope composition of the Alcáçovas calc-alkaline rocks (Ossa-Morena Zone, Portugal)

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The Alcáçovas area is located in the SW sector of the Ossa-Morena Zone (OMZ), close to a major fault that separates this geotectonic unit from the South Portuguese Zone (SPZ). Along this boundary, in the OMZ, testimonies of low-K tholeiitic and calc-alkaline magmatism are common and have been interpreted as being related to the operation of a subduction zone between OMZ and SPZ during the Variscan cycle [1]. Two main igneous lithologies, both displaying calc-alkaline compositions, can be found in the studied area: gabbro-diorites and dacitic-rhyolitic porphyries [2,3]. Outcrop conditions have not yet allowed to establish unequivocally the sequence of magma emplacement. In previous geochronological studies on the porphyries, whole-rock Rb-Sr dates and K-Ar ages cluster around 320 Ma [4,5,6].

According to field observations, sometimes felsic dykes cut mafic rocks, but there are also gradual transitions from gabbroic to tonalitic compositions, within bodies mapped as gabbro-diorite, revealing that different melts coexisted.

In this study, rock samples of both gabbro-dioritic bodies and porphyries were analysed for Rb-Sr and Sm-Nd isotopes. Considering the whole set of samples, no isochron was obtained, showing that they can not be simply related by crystal fractionation processes.

Rb-Sr data of porphyries from a single quarry (at Lameira, 7 km to the SW of Alcáçovas) give  $323 \pm 16$  Ma (MSWD=1.9; initial  $^{87}\text{Sr}/^{86}\text{Sr}=0.7097 \pm 0.0018$ ). Taking into account that the rocks of the Lameira outcrop show strong hydrothermal alteration, this date must be viewed as a consequence of a very efficient redistribution of mobile elements during aqueous fluid circulation and, as such, it places a minimum limit to the actual magmatic age

The plot of compositions of the gabbro-dioritic bodies, including their transitions to tonalites and the associated felsic dykes, in the  $\epsilon_{\text{Nd}}-^{87}\text{Sr}/^{86}\text{Sr}$  diagram, define an almost perfect hyperbole (from  $\epsilon_{\text{Nd}_{323}} = +3.9$  and  $^{87}\text{Sr}/^{86}\text{Sr}_{323} = 0.7058$  to  $\epsilon_{\text{Nd}_{323}} = -3.8$  and  $^{87}\text{Sr}/^{86}\text{Sr}_{323} = 0.7085$ ), as expected in a mixture between mantle-derived melts and crustal materials. In the same diagram, samples from the Lameira quarry show an almost constant  $\epsilon_{\text{Nd}_{323}}$ , between -2.4 and -2.9, and  $^{87}\text{Sr}/^{86}\text{Sr}_{323}$  varying from 0.7092 to 0.7106. Therefore, the Lameira porphyries could represent a member of the same mixture, with the Sr signature modified by hydrothermal fluids with a stronger crustal component.

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## Correlation of magnetic susceptibility with $\delta^{18}\text{O}$ data in magnetite- and ilmenite-type granites from Iberian massif

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The relationship between oxygen isotopic values and magnetic susceptibility composition on 11 Variscan Portuguese granites has been investigated. Whole-rock oxygen-isotope ( $\delta^{18}\text{O}$ ) values for Vieira do Minho (VM), Vila Pouca de Aguiar (VPA), Chaves, Castelo Branco (CB), Manteigas and Serra da Estrela (SE) granitoids, were compiled from bibliography [1,2,3,4], and  $\delta^{18}\text{O}$  for Santa Eulalia Plutonic Complex (SEPC) were obtained by laser fluorination at the Stable Isotopic Laboratory of Salamanca. Magnetic susceptibility (Km) values were obtained with a Kappabridge equipment from Toulouse University and Geology Centre, Porto University [2,5,6,7,8]. In this study is shown that there is a significant inverse correlation between Km and  $\delta^{18}\text{O}$ . *Magnetite*-type granites (Manteigas granodiorite and SEPC external facies) have  $\text{Km} > 10^{-3}$  SI and low  $\delta^{18}\text{O}$  values ranging from 8.9 to  $10.3\text{‰}$  instead those of *ilmenite*-type (all the other granites) have  $\text{Km} \leq 10^{-4}$  SI and are  $\delta^{18}\text{O}$  enriched (9.3 to  $13.5\text{‰}$ ). The *I*-type granites (VM, VPA, Chaves, Manteigas and SEPC external facies) show lower average  $\delta^{18}\text{O}$  ( $10.2\text{‰}$ ) and higher Km values ( $100 \times 10^{-6}$  SI) than the *S*-type granites (SE and CB) with  $\delta^{18}\text{O} = 12.6\text{‰}$  and  $\text{Km} = 65 \times 10^{-6}$  SI.

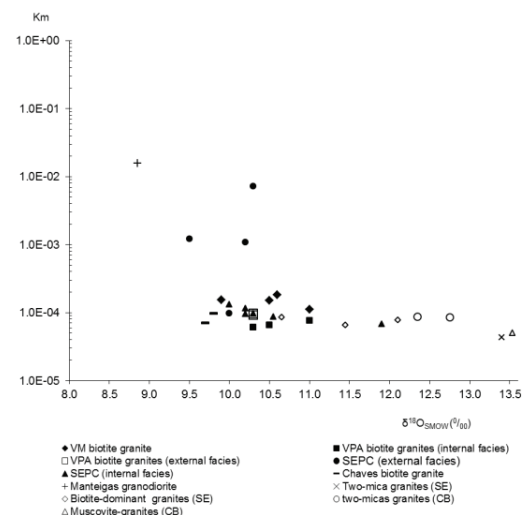


Figure 1: Semi-log plot of Km (in SI units) versus  $\delta^{18}\text{O}$ .

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