

Wireline Geophysical Logging of the Nirex Deep Boreholes in the Sellafield Area: Comparisons between BVG Core and Wireline Derived Formation Factors

A report produced for United Kingdom Nirex Ltd

Commissioned Report CR/02/168N

COMMISSIONED REPORT CR/02/168N

Wireline Geophysical Logging of the Nirex Deep Boreholes in the Sellafield Area: Comparisons between BVG Core and Wireline Derived Formation Factors

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Key words Sellafield; wireline logs; environmental corrections

Front cover

FMS (left) and BHTV (right) Images with scanned core photograph of fractured Borrowdale Volcanic Group rocks, Sellafield Borehole RCF1

Bibliographical reference

N R Brereton and P D Jackson. 2003. Wireline Geophysical Logging of the Nirex Deep Boreholes in the Sellafield Area: Comparisons between BVG Core and Wireline Derived Formation Factors. *British Geological Survey Commissioned Report*, CR/02/168N 38pp

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1. INTRODUCTION

The formation (resistivity) factor of a rock unit is a parameter based on electrical measurements that can be related directly to the porosity and, to a lesser degree, to the permeability. A reasonable correlation had previously been found between core sample and wireline derived formation factors for the Borrowdale Volcanic Group (BVG) in Nirex Borehole RCF3 in the Sellafield area (Brereton et al., 1996). It was concluded that wireline logs are able to provide an effective means of estimating the broad characteristics of formation factor variability with depth in a borehole and furthermore that they may be suitable for use in estimating the rock matrix diffusion properties of the *in situ* rock.

The objective of the work reported here has been to further test this approach using core sample and wireline derived formation factors from other Nirex deep boreholes in the Sellafield area. For the purposes of calculating the wireline derived formation factors, the method used in Brereton et al., 1996 has been adopted in which the pore water resistivity in the BVG in the vicinity of any particular borehole has been taken to be constant. Values of pore water resistivity have been estimated from measurements on ground water produced from the Nirex boreholes and on pore water extracted from core samples. The wireline derived formation factors have then been compared with formation factor measurements made by the BGS on core samples during the Nirex Core Characterisation Programme.

2. THE EVALUATION OF FORMATION FACTORS

The electrical properties of fluid saturated rocks are determined by the conductivity of the mineral grains, the conductivity of the pore fluids, and the rock porosity, which determines the relative effect of the two previous factors. Rock forming minerals are mostly silicates having very high resistivities (in the range 10^6 to 10^{14} Ω m) while the resistivities of natural groundwaters can range from less than 0.1Ω m to greater than 10 Ω m, depending upon the total dissolved solid concentrations.

The rock matrix rarely comprises an assemblage of grains of a single mineral species and most geologic minerals are a mixture of different materials, according to rock type. Diagenetic and mineralising processes will tend to redistribute the components of the existing rock matrix materials and produce new minerals that may locally be more concentrated. The direct influence of the rock matrix resistivity is generally relatively small and the wide range of measured bulk resistivities of water-saturated rocks primarily reflects the combined influence of the rock porosity and the associated pore fluid resistivity.

The formation factor, *F*, first introduced by Archie (1942), may be expressed as a ratio of the bulk resistivity of the saturated rock, *R*0, to the resistivity of the pore fluid *R*w:

$$
F = \frac{R_0}{R_w} \tag{2.1}
$$

Archie (1942) showed that the empirical relationship between F and the porosity ϕ is of the form

$$
F = \frac{1}{\phi^m} \tag{2.2}
$$

where m is the cementation factor whose value lies between 1.3 and 2.5.

Provided the mineral phase can be assumed to be an insulator and the pore water is sufficiently saline (to limit surface conduction) the formation factor can be regarded as an intrinsic parameter or the rock, related only to the geometry of the transport porosity (Brereton et al., 1996).

3. FORMATION FACTORS FROM CORE SAMPLES

Special precautions need to be taken to preserve the *in situ* fluids present in the core during drilling operations and to prevent the rock cores from subsequently drying out. Worthington *et al.* (1988) applied Archie's equation to both preserved and resaturated sedimentary core samples with porosities of about 20 %. They attributed differences in the results to induced changes in pore geometry, and thence to surface conduction, through the irreversible collapse of structurally delicate clay minerals related to the passage of a fluid interface through the sample during drying of the core plugs. Although this example may not be directly relevant to low porosity volcanic rocks such as the BVG, it serves to demonstrate the need for caution when preparing cores for laboratory evaluation. No such special precautions were taken with regard to the core samples used in this study.

In the absence of these special precautions, the core samples will generally be resaturated with a fluid of known properties. This procedure will render it unrealistic to make direct comparisons between wireline and core sample resistivity measurements, but formation factors, which represent the dimensionless ratio between the bulk resistivity of the saturated rock and the pore fluid resistivity, are more directly comparable.

As part of the Nirex Core Characterisation Programme on the drill core from the Nirex deep boreholes, the BGS carried out a total of 316 BVG resistivity measurements on core samples from Boreholes 2, 3, 4, 5 and 7A (Nirex, 1997a). Core plugs, having dimensions of 38 mm diameter by 76 mm long, were taken from the core sticks using standard cutting techniques. No special precautions were taken with regard to the fluids used during this cutting procedure and so some crosscontamination between the cutting fluids and the *in situ* pore fluids may have occurred. The core plugs were then vacuum dried and resaturated using fluids with compositions intended to simulate BVG groundwaters. Different fluid compositions were used at different depths in a given borehole using information provided from other Nirex programmes at the time. The resaturating fluid resistivities ranged from 0.16 to 0.30 Ω m. Resaturation was allowed to occur under vacuum for a period of several hours prior to resistivity measurements being made. The measurement temperatures ranged from 18 to 23 °C.

No attempt was made to flush out any residual pore fluids prior to resaturation; or salts that may have deposited while the core sticks were drying out; or any residual cross-contamination from fluid that may have entered during core plug cutting. Nor was any attempt made to establish whether the interstitial pore fluids had fully reached chemical equilibrium during the resaturation procedure. However, given the very low interstitial porosity of the BVG and the relatively large size of the core plug samples, it is unlikely that cutting fluid contamination would have been too significant. Also, because the resaturating fluids were intended to be similar in composition to the BVG groundwaters, the pore fluids experienced during the bulk resistivity measurements would be expected to exhibit similar electrical properties to the original *in situ* pore fluids. Nevertheless, uncertainties will remain in these expectations.

The measured saturated sample resistivities, water resistivities and formation factors (*Fc*) derived from core samples are listed in Table 3.1.

4. BVG FORMATION FACTORS FROM WIRELINE LOGS

4.1 Wireline measurements of formation resistivity

Boreholes are usually drilled with fluids other than the formation water and may either be fresh or saline water. Depending upon the local formation permeability and porosity distributions, and also upon relative difference between the *in situ* formation pore pressure and the borehole drilling fluid pressure, these drilling fluids will permeate into the rock formation. To accommodate these varying conditions near the borehole wall, a range of logging tools have been developed by the logging service companies, each with differing capabilities designed to penetrate greater or lesser distances into the formation (deep, medium or shallow) and also to provide greater or lesser focusing for improved vertical resolution. Induction tools $(\mathbf{ILD}^{\mathbb{O}})$ and $\mathbf{ILM}^{\mathbb{O}})$ are usually used where the formation is more electrically conductive than the borehole fluid and focused laterolog tools (LLD^{\circledcirc}) and $LLS^{\circledcirc})$ in other situations. The Nirex boreholes were predominantly logged using the LLD[©] and LLS[©] tools.

In low permeability, low porosity and relatively homogeneous rocks, where drilling fluid invasion is minimal, it may be expected that the deep and shallow resistivity measurements will be similar because they will both sample similar pore fluid characteristics. This is borne out by the borehole wireline log resistivity profiles, where the deep and shallow logs from the BVG tend to follow one another very closely. In Brereton et al., 1996 it was concluded that the deep resistivity LLD log is better able to represent the relationship between resistivity and porosity, as described by the Archie equation (2.2), than the shallow resistivity LLS log. Therefore the LLD resistivity log was previously used to evaluate an *in situ* formation factor log for Borehole RCF3 (Brereton et al., 1996). This practice has been adopted here.

4.2 Evaluation of *in situ* **fluid resistivities**

If quantitative formation factor assessments are to be made from the wireline logs, then estimates of the *in situ* pore fluid resistivity within the formations around each borehole are required. As part of the Nirex hydrogeological investigations, a series of borehole hydraulic tests were performed during which pumped water samples were collected from selected depth intervals. Chemical analyses were carried out on these water samples and, in many cases, electrical conductivity measurements were also made.

To characterise the chemistry of the interstitial pore fluids, experiments were conducted by both the BGS and AEA Technology using Boreholes 2 and 3, to extract pore waters from rock cores by leaching and by centrifugation (Nirex, 1992, Report No. 202 and Nirex, 1993, Report 213). Ranges of chemical constituents were analysed for, but neither fluid conductivities nor resistivities were included in those measurements. To overcome this, a correlation between chloride concentration and electrical conductivity was constructed from the water sample data collected during the hydraulic testing programme (Brereton et al., 1996). This took the following form:

$$
\sigma_f = 0.1486 + 0.238 \, 10^{-3} \, Cl - 0.486 \, 10^{-9} \, Cl^2 \tag{4.1}
$$

Where σ_f is the electrical conductivity of the fluid (S m⁻¹),

Cl is the chloride concentration (g m⁻³).

All available hydraulic test interval and pore water sample data from the BVG sections of Boreholes 2, 3, 4, 5 and 7A were collated and Equation 4.1 was used to convert measured chloride concentrations into electrical conductivity and then into fluid resistivity. The results are listed in Table 4.1. [Note: conductivity is tabulated as ms/cm where y mS/cm is equivalent to $10/\text{y} \Omega$ m

Table 4.1 also gives the mean fluid resistivity value for each borehole. Some individual values were excluded from these means for the reasons given below. Also, where hydraulic test interval fluid resistivity values have been derived both directly from conductivity measurements and indirectly from chloride measurements (via Equation 4.1), then only the conductivity-derived values have been used in the mean. This is because although the conductivity and chloride are independent measurements, the chloride-derived resistivity is based on a correlation between these two, and so, it might be argued, double counting may occur.

For Borehole 3, four of the interstitial pore water sample resistivity values (out of 22) were significantly greater than the rest (averaging about 0.61 Ω m) and so were considered anomalous and excluded from the mean. Similarly, for Borehole 7A two of the interstitial pore water sample resistivity values (out of 9) were much greater than the rest (averaging about 0.28 Ω m) and these were also excluded from the mean. All the pore water data from Borehole 4 are considered to be anomalous (Steve Swanton, AEAT; personal communication), and have also been excluded from the mean. It is unclear whether or not these anomalous interstitial pore water sample values from Boreholes 3, 4 and 7A can be attributed to sample preparation and handling or to genuine geological reasons.

Allowing for these exclusions, the resistivities of the interstitial pore water samples are, in general, very similar to those of the hydraulic test interval water samples. However, there is a tendency for the pore water sample resistivities to be slightly higher than the pumped water sample resistivities. This implies that the fluids

flowing through the network of fractures are marginally more saline than the interstitial pore waters within the body of the rock.

The mean values of the pumped water samples for Boreholes 2, 3, 4, 5 and 7A are $0.27, 0.05, 0.30, 0.27$ (one sample) and 0.11Ω m. The mean values of the interstitial pore water samples for Boreholes 2, 3 and 7A are 0.39, 0.07 and 0.15 Ω m (Table 4.1). For comparison, the mean pumped water and pore water sample resistivities from below the Saline Transition Zone for all lithologies from all the Nirex boreholes (representing 122 hydraulic test intervals and 183 pore water samples) are 0.23 ± 0.18 and 0.22 ± 0.23 Ω m respectively (Brereton et al., 1996).

Water column fluid conductivity wireline logs were recorded in some boreholes during the water abstraction tests. These were primarily for the purpose of identifying fluid flow horizons under flowing conditions. Repeat water column logging runs were carried out over several days in Boreholes 4 (three runs) and 5 (four runs), while in Borehole 2 and 7A single profiles were recorded (Figure 4.1). All these profiles exhibit a gradational decrease in fluid resistivity with increasing depth. In Borehole 2 the profile decreases from 0.27 to 0.16 Ω m in a series of stepwise changes. The three logging runs in Borehole 4 decreased from about 0.70 to 0.28 Ω m and also decreased with time in the upper part of the profile. In Borehole 5, the first logging run showed a fairly linear gradient, decreasing from about 0.79 to 0.59 Ω m. Subsequent logging runs showed distinctive profiles, superimposed upon the linear gradient, which deviate away from the gradient line towards successively lower fluid resistivity values at specific borehole depths. These indicate distinct fluid flow horizons into or out of the borehole. The fluid resistivities at these depth horizons successively trend towards a value of about 0.45 Ω m or even lower. The single profile in Borehole 7A decreases from 0.80 to 0.47 Ω m and is similar in character to those in Boreholes 4 and 5.

Fluid resistivity is influenced by temperature. Quist and Marshall (1969) demonstrated significant increases in electrical conductivity (and hence decreases in fluid resistivity) of sodium chloride solutions up to temperatures of 400ºC. They also demonstrated that increases in pressure of up to 0.4 GPa over this temperature range had little influence.

Much of the gradational decreases in fluid column resistivity with depth in the Nirex boreholes can be attributed to temperature changes with depth. In Borehole 2 for example, the wireline logs show that from 475 to 1590 mbRT the temperature $(T; {}^{\circ}C)$ increases with depth $(D; m)$ from 19.3°C to 47.2°C according to the following relationship:

$$
T = 0.025D + 7.4
$$

Over the same depth range a near linear relationship between fluid resistivity (R_w) and temperature approximates to:

$$
R_w = -0.0038T + 0.33\tag{4.3}
$$

Therefore, at a temperature of say 20ºC, which is within the range at which most of the measurements on the pumped water samples were made, the fluid resistivity

would be 0.25 Ω m, which is similar to the mean value of 0.27 Ω m directly measured on the pumped water samples from this borehole (Table 4.1).

These fluid resistivity profiles provide only limited quantitative information. Because of vertical fluid movements within the borehole water column, they will be more representative of fluids entering the borehole from higher permeability zones at different depths than of the interstitial pore fluids within the formation at the depth of measurement. Equally, fluid resistivities of pumped water samples, from relatively narrow intervals isolated by packers, will be more representative of fluids flowing through discrete fractures rather than of the interstitial waters, which may explain the tendency for the pore water sample resistivities to be slightly different to the pumped water sample resistivities.

Even so, it would be expected that there would be an overall agreement between the general ranges of values. Allowing for the various caveats outlined above, and for the effects of temperature on the fluid column wireline resistivities, there does appear to be a general agreement. Therefore, it appears to be reasonably justified to assume a constant *R*w for each borehole, as was concluded by Brereton et al., 1996.

There is a further aspect that needs consideration, the effects of temperature. The wireline measurements of formation resistivity at any given depth in the borehole are made at the prevailing *in situ* temperature at that depth. These formation resistivity measurements are not subsequently adjusted to the ambient temperature at ground level. For each borehole, the derivation of wireline formation factors based upon Equation 2.1 should, in principle, incorporate a correction to R_w for temperature changes with depth similar to Equations 4.2 and 4.3. Wireline temperature logs were not run in all these boreholes and even where they were they would be subject to perturbations due to fluid movements within the water column similar to those that affect the fluid conductivity logs described above (Figure 4.1). Because of these difficulties temperature corrections to R_w have not been attempted.

Therefore, for the purposes of making comparisons between the core and wireline derived formation factors, a constant R_w has been assumed for the BVG around each borehole. It should be recognised that due to the effects of temperature changes with depth and the other factors discussed, the actual three-dimensional distribution of fluid resistivity within the BVG will be variable.

For simplicity, the mean R_w value derived from the pumped water samples and the interstitial pore water samples for each borehole (given the exclusions described previously), has been used as a basis for calculating formation factors from the wireline logs. No account has been taken of the borehole fluid column resistivity profile data for this purpose. For Boreholes 2, 3, 4, 5 and 7A these mean R_w values are 0.34, 0.06, 0.30, 0.27 and 0.14 Ω m respectively (see Table 4.1, means highlighted in bold).

5. DISCUSSION

5.1 Borehole / Core Comparison

Wireline derived formation factors (F_w) were calculated using Equation 2.1 where R_0 is the wireline deep resistivity (LLD) and R_w is the corresponding mean fluid resistivity. Cumulative frequency distributions for Boreholes 2, 3, 4, 5 and 7A are shown in Figures 5.1 to 5.5. Wireline derived formation factor profiles are shown in Figure 5.6 (black profiles), alongside the core sample formation factors (blue dots).

With the exception of Borehole 2 the core sample values are consistently less than the corresponding wireline derived values. It is clear from Figures 5.1 to 5.6 that in some cases the differences can be large and range to more than one order of magnitude. The largest differences occur in the upper section of Borehole 4 (above about 830 m; Figure 5.6) and the lower section of Borehole 5 (below about 1170 m). In the case of Borehole 4, this upper section coincides with the zone where significant departure from the general trend of the fluid resistivity profile is observed (Figure 4.1). This implies that it may be unreasonable to adopt the same mean R_w value of 0.30 Ω m over this depth zone as had been adopted for the whole borehole. Despite these offsets, the broad core sample value trends tend to follow the same broad trends followed by the wireline log profiles. This is particularly so for Boreholes 2, 3 and 7A.

For Borehole 2 the match is much closer. A correlation plot between core sample formation factors and wireline interpolated values for Borehole 2 is shown in Figure 5.7. The correlation is moderate, but this plot needs to be treated with caution owing to potential problems associated with the depth registration between the core and wireline data (Nirex Report SA/97/021). Although all core depths were corrected to the wireline log depths as part of the Core Characterisation Programme, errors will remain where core loss has occurred. These problems are exacerbated when dealing with logarithmic data in that small core to wireline depth misalignments could result in a large difference between formation factors. That is, the Figure 5.7 correlation might be better if absolute depth matching between the core and wireline data could be guaranteed.

Figure 5.8 demonstrates that a closer match between the core and wireline derived Formation Factors can be derived by using a fixed value of 035 ohm metres for all the boreholes.

Summary statistics for the core sample formation factors listed in Table 3.1 are given in Table 5.1a. It is noteworthy that the core sample formation factors for Borehole 2 are very much higher and more varied than those for the other boreholes, while those for Borehole 3 are relatively low. Summary statistics for the wireline derived formation factors are given in Table 5.1b. It is clear that not only are the ranges of wireline derived formation factors much greater than the core sample values but the mean values are also much greater.

5.2 General Review

It is clear from these comparisons that although the core sample formation factors broadly follow the variations in the wireline log derived profiles, there appear, with some exceptions, to be systematic differences, with the core sample values generally being less than the wireline derived values.

Formation factors are determined as a simple dimensionless ratio between the measured saturated bulk rock resistivity and the fluid resistivity. The wireline methods for measuring resistivity in boreholes are well established and well understood (Desbrandes, 1985) and there is little reason to question the wireline resistivity results. There is a broad agreement between the general ranges of formation fluid resistivity values, but some uncertainties remain because of the relatively small *in situ* fluid resistivity sample data sets and because of systematic differences between the pumped water and interstitial pore water resistivities. There will, therefore, be corresponding uncertainties in the wireline derived formation factors.

The core samples were saturated with fluids of known resistivity prior to the measurement of bulk resistivity. However, there are uncertainties as to whether, during this procedure, the interstitial pore fluids had fully reached chemical (thus electrical) equilibrium with the resaturating fluids. These uncertainties will carry through to the core sample formation factors.

The main uncertainties in making comparisons between core and wireline derived formation factors lie with those associated with the resaturation of the core samples and with the establishment of best estimates of fluid resistivity for the derivation of wireline formation factors.

It was stated in Section 3 that no deliberate attempt was made to flush out any residual pore fluids prior to resaturation of the core samples, or salts that may have deposited while the core sticks were drying out. On resaturation this may lead to the core samples containing water that is significantly more saline than the resaturating solution. For example, following resaturation of the Borehole 3 core samples, the actual pore fluid resistivity may have been closer to the value used to calculate the wireline derived formation factors (i.e. 0.06 Ω m), than the resistivity of the core sample resaturating solution (i.e. 0.23 Ω m). If this were true, then the mean core sample formation factor of 303 given in Table 5.1a would be about 1,200, which is much closer to the mean wireline derived formation factor of 1,398 given in Table 5.1b.

It is worth noting that, under the Nirex Safety Assessment Research programme, Borehole RCF3 core samples were passively resaturated for 3 to 5 months prior to resistivity measurements. Good agreement was found between formation factors measured after 3 and 5 months (Brereton et al., 1996) indicating that electrical equilibrium had been reached. For these RCF3 samples a better correlation between the core and wireline formation factor data was found than for the five boreholes considered here.

An additional possibility, that may partially explain the differences between the wireline and core derived formation factors, is that, during the coring and sub-coring operations, de-stressing of the core samples will have taken place which may result in slight increases in porosity. This, in turn, would result in a decrease in the measured core sample formation factors relative to what would have been measured had destressing not occurred. Since the core porosity is typically between 0 and 5 per cent in the BVG small per cent change in porosity could have dramatic effects on the Formation Factor.

Also, the effects of core sample bias during sampling can be significant. The core will for example only be sampled where it is relatively intact and also there will be a bias to sample a variety of rock types rather than sample at random intervals. The differences between core and wireline scales of measurement can also have a significant impact (Brereton et al., 1996) and are beyond the scope of this report.

6 CONCLUSIONS

This technical note describes the results of comparisons between BVG core sample and wireline derived formation factors for five of the Sellafield boreholes. The objective of these comparisons was to further test the reasonable correlation found previously in Borehole RCF3 (Brereton et al., 1996).

Although the resistivities of the interstitial pore water samples for the five boreholes are very similar to those of the hydraulic test interval water samples, there is a tendency for the pore water sample resistivities to be slightly higher than the pumped water sample resistivities. This implies that the fluids flowing through the network of fractures are marginally more saline than the interstitial pore waters within the body of the rock. Nevertheless, for the purposes of making comparisons between the core and wireline derived formation factors, it appears to be reasonably justified to assume a constant *R*w for the BVG around each borehole, as was concluded by Brereton et al., 1996. However, no attempt has been made here to correct R_w for formation temperature changes with depth.

Two principal conclusions arise out of the comparisons between the core and wireline derived formation factors. Firstly, the core sample values broadly follow the variations in the wireline log derived profiles. Secondly, while for Borehole 2 there is a reasonable quantitative agreement, for Boreholes 3, 4, 5 and 7A there are systematic differences, with the core sample values being generally less than the wireline derived values.

It is very likely that these differences may largely be attributed to uncertainties associated with the question as to whether or not the core sample interstitial pore fluids had reached equilibrium during the resaturation process. Also, to a lesser degree, to de-stressing of the core samples during the coring and sub-coring operations; to core sample bias and scale of measurement effects; and with the establishment of best estimates of the *in situ* formation fluid resistivities.

In general, it may be concluded that, because of the additional uncertainties about the equilibration of the resaturated core samples, the resistivity measurements made as part of the Core Characterisation Programme were not as good a test of the correlation between core and wireline formation factors as had been expected. To reduce this uncertainty, it is recommended that, in future work, the primary objectives of the core measurement programmes are clearly defined and that sufficient care is taken during core sample preparation to ensure that those objectives can be met. In the particular case of formation factor measurements, it is important to ensure equilibration of the core samples prior to resistivity measurements being made. On the basis of the present data, it is probable that for these five boreholes the wireline derived formation factors are more indicative of the *in situ* formation factor values in the field than are the core sample measurements reported here.

Given these caveats, the findings of this report support the conclusions of Brereton et al., 1996 in that wireline logs are able to provide an effective means of estimating the broad characteristics of formation factor variability with depth in a borehole.

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Table 3.1: Resistivity measurements on BVG core samples and derived formation factors (data abstracted from Nirex Report SA/97/017).

Table 4.1: BVG fluid resistivity from hydraulic test water sample electrical conductivity and chloride measurements and from pore water sample chloride
extraction measurements. NB: a) the overall means are for all the resi **Table 4.1: BVG fluid resistivity from hydraulic test water sample electrical conductivity and chloride measurements and from pore water sample chloride extraction measurements. NB: a) the overall means are for all the resistivity data for each borehole, but exclude those values in underlined italics. b) data sources are: (1) BGS (Nirex Report S/96/001; Nirex Report Nos. 202 and 203): leachate; (2) Nirex Report SA/95/008; (3) Nirex Report SA/97/089; (4)AEA: leachate** (average of shape and density determinations). **(average of shape and density determinations).**

Comparisons between BVG Core and Wireline Derived Formation Factors Comparisons between BVG Core and Wireline Derived Formation Factors

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Table 5.1: Summary statistics for BVG formation factor values.

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Borehole	BH2	BH ₃	BH4	BH ₅	BH7A	
Mean	2,587	303	684	613	560	
Standard deviation	2,066	115	272	309	336	
Maximum	12,404	591	1,389	1,685	1,250	
Minimum	245	109	199	136	105	
Number	97	35	80	75	29	

(a) Core sample formation factors (*Fc***); blue dots in Figure 5.6**

Figure 4.1: Fluid resistivity profiles derived from wireline conductivity logs through the water column of Boreholes 2, 4, and 5 over the BVG depth range (fluid resistivities in Ω **m; depths in mbRT).**

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Figure 5.6: Formation factor profiles for Boreholes 2, 3, 4, 5 and 7A over the BVG depth range (formation factors are on a logarithmic scale from 10 to 90,000; depths in mbRT). The black profiles are wireline derived values using the fluid resistivities described in the text. The blue dots are core derived values.

Figure 5.8: Formation factor profiles for Boreholes 2, 3, 4, 5 and 7A over the BVG depth range (formation factors are on a logarithmic scale from 10 to 90,000; depths in mbRT). The green profiles are wireline derived values using a single fluid resistivity of 0.35 Ω m for all boreholes. The blue dots are core derived values.