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北山东南部早白垩世伸展构造变形：二维反射地震剖面解释与磷灰石裂变径迹测年的制约

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The Early Cretaceous extensional deformation in the southeastern Beishan Range, central Asia: Constrains from 2D seismic reflection profile interpretation and apatite fission track thermochronology

Abstract: [Objective] The Beishan Range occupies a key position in Central Asia. This study aims to deepen the understanding of the timing, intracontinental deformation processes, and their dynamic mechanisms in the Late Mesozoic on the southern margin of the Central Asian Orogenic Belt (CAOB). [Methods] We conducted detailed analyses of the Early Cretaceous extensional and earlier compressional structures in the southeastern Beishan Range through field geological observations, interpretation of 2D reflection seismic profiles, and apatite fission track thermochronology. [Conclusion] Field observations show that Lower–Middle Jurassic strata have been strongly deformed by numerous thrusts and folds. 2D seismic reflection profiles reveal two NE- to NEE-striking normal faults. The Suosuojing fault is a SE-dipping low-angle listric normal fault, while the Wudaoming fault is a NW-dipping high-angle normal fault. These normal faults cut through the early-formed fold-thrust system, indicating a transition from contraction to extension. The Suosuojing and Wudaoming faults, respectively, define the northwestern and southeastern boundaries of the Early Cretaceous Zongkouzi basin. The Zongkouzi basin exhibits a graben geometry, with Lower Cretaceous strata displaying typical growth-strata relationships, suggesting that the normal faults were active during the late Early Cretaceous. Thermal history modeling results from apatite fission track data indicate that the footwall of the Suosuojing fault experienced rapid cooling between 132 and 110 Ma. This rapid cooling phase was closely related to the footwall exhumation during the normal slip of the Suosuojing fault. We argue that the Late Mesozoic intracontinental contraction–extension transition in the southeastern Beishan Range likely occurred between ~133 Ma and ~129 Ma in the late Early Cretaceous. The collapse of the thickened crust and coupled mantle upwelling triggered the Early Cretaceous extensional deformation in the southern CAOB.

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摘要: 为深入认识中亚造山带南缘晚中生代陆内变形过程及其动力学机制,通过野外地质观察、二维反射地震剖面解释及磷灰石裂变径迹测年,对北山东南部早白垩世伸展构造及早期挤压构造进行了详细解析。结果表明,一系列逆冲断层与褶皱构造造成下—中侏罗统发生强烈的挤压变形。地震剖面揭示出2条早白垩世伸展正断层,其中梭梭井断层为南东倾向的低角度铲式正断层,五道明断层为北西倾向的高角度正断层,二者共同切割了早期形成的褶皱–冲断系统,指示挤压–伸展构造的转换;梭梭井断层与五道明断层分别限定了早白垩世总口子盆地的北西和南东边界,使得其具有“地堑”样式,盆地内沉积的下白垩统生长地层发育,表明伸展正断层的活动时间为早白垩世晚期。磷灰石裂变径迹热史模拟结果显示,梭梭井断层下盘于132~110 Ma经历了快速冷却和剥露事件,该事件与其持续的正断层活动密切相关,进一步证实北山东南部晚中生代挤压–伸展构造的转换很可能发生在早白垩世晚期(133~129 Ma)。增厚地壳的重力垮塌与局部地幔上涌共同导致了中亚造山带南缘早白垩世的区域伸展作用。

关键词: 北山; 早白垩世; 伸展构造变形; 磷灰石裂变径迹; 挤压–伸展构造转换

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0 引言

中亚地区发育众多典型的陆内变形的实例,如新生代的天山、阿勒泰及祁连山等(Yin et al., 1998; Cunningham, 2005; Jolivet et al., 2010; Zuza et al., 2018)。在新生代印度–亚洲板块碰撞之前,中亚地区中生代经历了极其复杂的变形过程(Zheng et al., 1996; Dumitru and Hendrix, 2001; Jolivet et al., 2001; Darby and Ritts, 2002; Wang et al., 2022),记录了亚洲陆缘不同块体之间多期汇聚与碰撞拼贴的过程(Darby and Ritts, 2002; Dong et al., 2015; 董树文等, 2019; Chen et al., 2022; Liu et al., 2023)。另外,中生代复杂的变形模式很有可能影响到了中亚地区新生代的构造变形分布(Dumitru et al., 2001; Jolivet et al., 2001, 2010; Zuza et al., 2018; Cheng et al., 2019)。因此,中亚地区成为研究陆内变形过程和探讨变形机制的热点地区(Cunningham, 2005; Raimondo et al., 2014)。

北山位于中亚造山带的最南缘,同时紧邻特提斯–喜马拉雅–青藏高原造山带的最北缘,构造位置十分关键(Wang et al., 2010; Xiao et al., 2010; Li et al., 2023)。已有研究表明,晚中生代陆内挤压变形将古生代一早中生代北山造山带构造活化,形成了北山褶皱–冲断带(左国朝等, 1992; Zheng et al., 1996; Dumitru and Hendrix, 2001; Liu et al., 2023)。北山褶

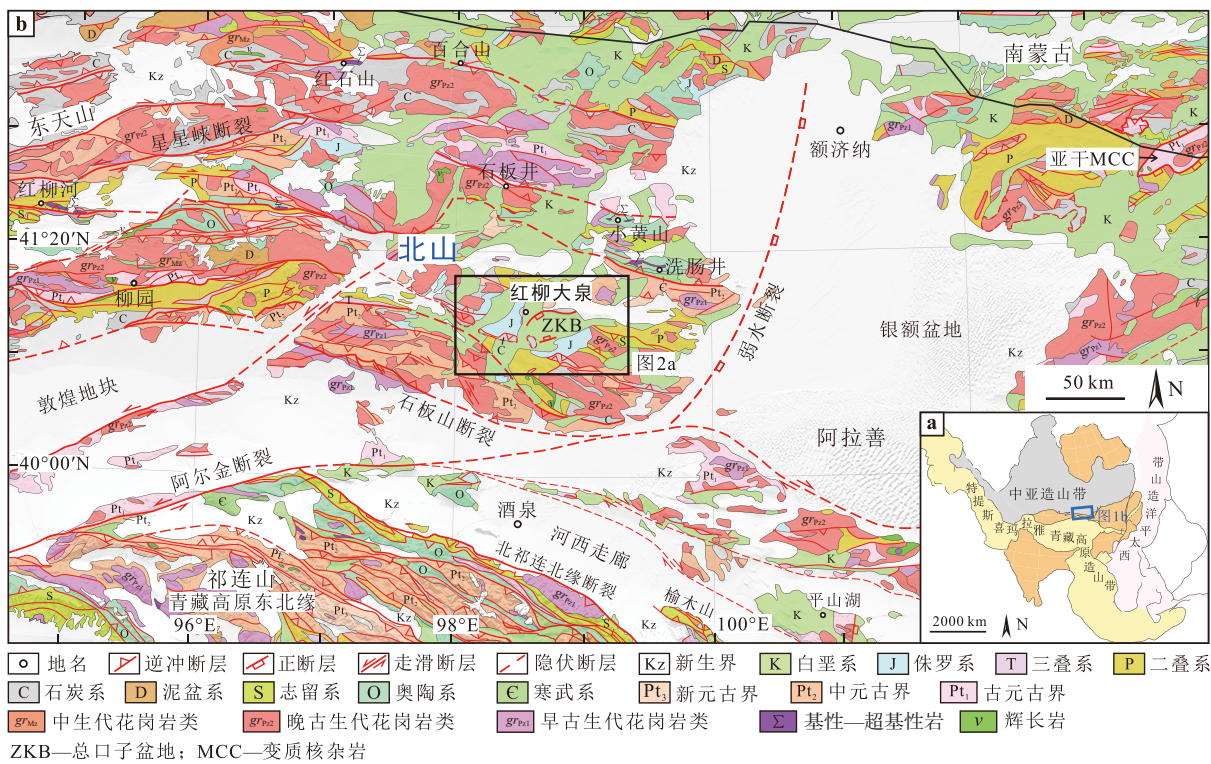
皱–冲断带开始形成于中侏罗世晚期,挤压逆冲作用可能一直持续到早白垩世。同时,热年代学数据也记录了早白垩世晚期(124~115 Ma)与正断层活动相关的下盘剥露与冷却事件,指示晚中生代北山褶皱–冲断带经历了挤压–伸展构造的转换(Liu et al., 2023)。区域上,这一构造转换普遍存在于中亚造山带南缘,形成了众多叠加于褶皱–冲断带或增厚地壳之上的正断层、断陷盆地及变质核杂岩等(Webb et al., 1999; Johnson et al., 2001; Chen et al., 2003; Meng, 2003; Cunningham et al., 2009; Wang et al., 2011; 朱日祥等, 2012; Lin et al., 2013; Zhu et al., 2015)。截至目前,部分学者认为北山、南蒙古及阿拉善地区伸展构造变形开始于晚侏罗世(Zheng et al., 1996; Graham et al., 2001; Meng, 2003)。然而,诸多研究结果表明区域伸展作用开始于早白垩世(Davis et al., 2002; Davis and Darby, 2010; 张长厚等, 2011; Wang et al., 2018; Lin et al., 2019, 2020; Liu et al., 2023)。同时,相关学者提出了多种关于中亚造山带南缘晚中生代陆内伸展构造变形的动力学模式,包括增厚地壳的重力垮塌(Meng, 2003; Wang et al., 2011)、古太平洋板块俯冲后撤引发的弧后伸展(朱日祥等, 2011, 2012; Zuo et al., 2020; Hui et al., 2021)及岩浆底侵(张宇等, 2024)等。但是,在北山地区仍缺少针对该期伸展构造变形样式、模式及起始时间等的研究。这些争议和不足制约了对中亚造山带南缘晚中生代陆内变形过程及其动力学机

制的认识。

为解决上述问题, 文章选择北山东南部红柳大泉地区作为研究区。基于卫星影像解译与野外地质观察, 分析了下一中侏罗统和下白垩统的构造变形特征; 并通过开展二维反射地震剖面解释, 详细解析了研究区内的伸展构造变形特征; 此外, 根据磷灰石裂变径迹测年及热史模拟结果, 进一步限定了伸展构造变形的时间。这些研究对深入认识中亚造山带南缘晚中生代陆内挤压-伸展构造转换及其动力学模式具有重要作用。

1 区域地质背景

北山位于中亚造山带的最南缘(图 1a; Zuo et al., 1991; Xiao et al., 2010; Li et al., 2023; 唐卫东等, 2023), 呈近东西向展布, 北接南蒙古构造带, 往西以星星峡断裂为界与东天山相接, 南邻敦煌地块与河西走廊, 往东与阿拉善地块相接(图 1b; Zuo et al., 1991; Wang et al., 2010; Xiao et al., 2010; Feng et al., 2018; Zhang et al., 2021a, 2022; 贺昕宇等, 2022)。



a—亚洲大地构造简图; b—北山地区及周缘区域地质图

图 1 亚洲大地构造简图与北山地区及其周缘区域地质图 (任纪舜等, 2013; 徐学义等, 2015; 陈宣华等, 2019; Liu et al., 2023)

Fig. 1 Sketched tectonic map of Asia and regional geologic map of the Beishan Range and its surrounding belts, central Asia (modified from Ren et al., 2013; Xu et al., 2015; Chen et al., 2019; Liu et al., 2023)

(a) Sketched tectonic map of Asia; (b) Regional geologic map of the Beishan Range and its surrounding belts

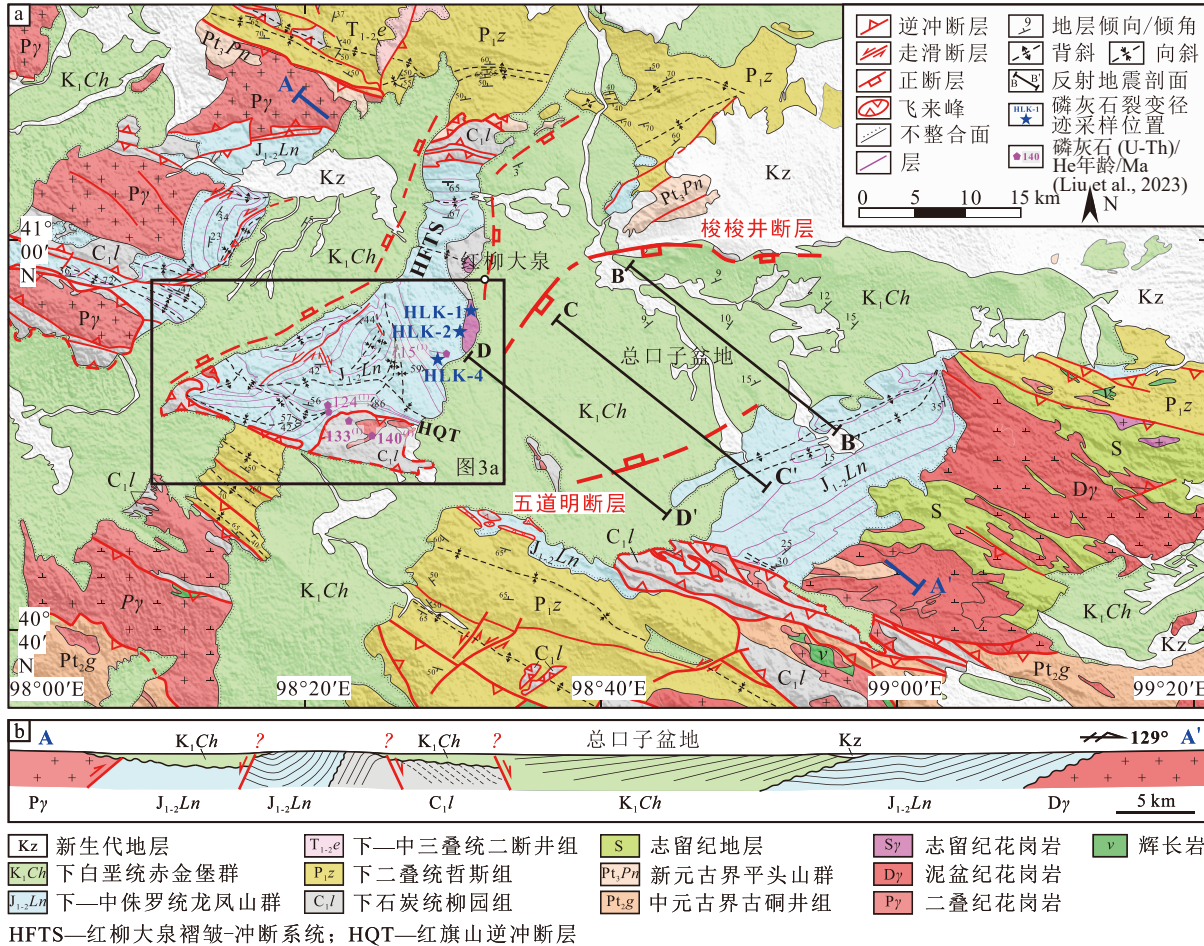
ZKB—Zongkouzi Basin; MCC—metamorphic core complex

北山地区出露前寒武纪地层, 古生代与俯冲、裂谷及碰撞造山作用相关的岩浆岩和变质岩, 古生代海相沉积地层, 以及中-新生代陆相沉积地层 (甘肃省地质局和第二区域地质调查队, 1971; Xiao et al., 2010; Song et al., 2013; Cleven et al., 2018; He et al., 2018; Li et al., 2023)。北山东南部下石炭统柳园

组 (C₁l) 主要由流纹斑岩、安山岩、凝灰质砂岩、砂岩和灰岩组成 (甘肃省地质局和第二区域地质调查队, 1971; 靳拥护等, 2020); 下二叠统哲斯组 (P₁z) 岩性为玄武岩、安山岩、凝灰岩、砾岩、砂岩及灰岩等 (甘肃省地质局和第二区域地质调查队, 1971); 下一中侏罗统龙凤山群 (J₁₋₂Ln) 由底部厚层砾岩和中一

上部的含砾粗砂岩、砂岩、粉砂岩及粉砂质泥岩组成(甘肃省地质局和第二区域地质调查队, 1971; 牛海青等, 2021); 下白垩统赤金堡群(K_1Ch)分布极为广泛(图 1b, 图 2a), 主要由砾岩、砂岩、泥质粉砂岩

和页岩组成(Li et al., 2007; 彭楠等, 2013; 张金龙等, 2017; 任文秀等, 2022)。新生代地层厚度很薄, 为一套冲积扇与河流相沉积(甘肃省地质局和第二区域地质调查队, 1971)。



a—红柳大泉地区地质图; b—北西—南东走向 AA' 剖面图

图 2 北山东南部红柳大泉地区地质图(底图据甘肃省地质局和第二区域地质调查队, 1971; 靳拥护等, 2020 修改)

Fig. 2 Geologic map of the Hongliudaquan area, southeastern Beishan Range (base map modified from BGGP, 1971; Jin et al., 2020; Apatite (U-Th)/He ages according to Liu et al., 2023)

(a) Geological map of the Hongliudaquan area; (b) NW-SE trending section AA'

HFTS—Hongliudaquan Fold-Thrust System; HQT—Hongqishan Thrust Fault

北山造山带形成于古生代—早中生代时期古亚洲洋多个分支洋盆长期俯冲消减、增生及最终闭合的过程中(Xiao et al., 2010; Ao et al., 2012; Song et al., 2013; He et al., 2018; Li et al., 2023)。随后, 强烈的陆内挤压变形将其构造活化, 形成了晚中生代北山褶皱—冲断带。其中, 北山逆冲推覆构造将大量中元古界白云质灰岩和大理岩等往北逆冲推覆至新元古界至下—中侏罗统之上, 其推覆距离至少约

为 120 km(左国朝等, 1992; Zheng et al., 1996)。在北山东南部, 红柳大泉褶皱—冲断系统将下—中侏罗统强烈挤压变形(图 2a; Liu et al., 2023)。该挤压变形系统沿南北和北西—南东 2 个方向同时发生水平缩短(收缩变形)。红柳大泉褶皱—冲断系统被下白垩统不整合覆盖(图 2)。区域上, 早白垩世地层普遍被认为沉积在一系列地堑/半地堑盆地中(Meng, 2003; Meng et al., 2003; 张金龙等, 2017; Hui et al.,

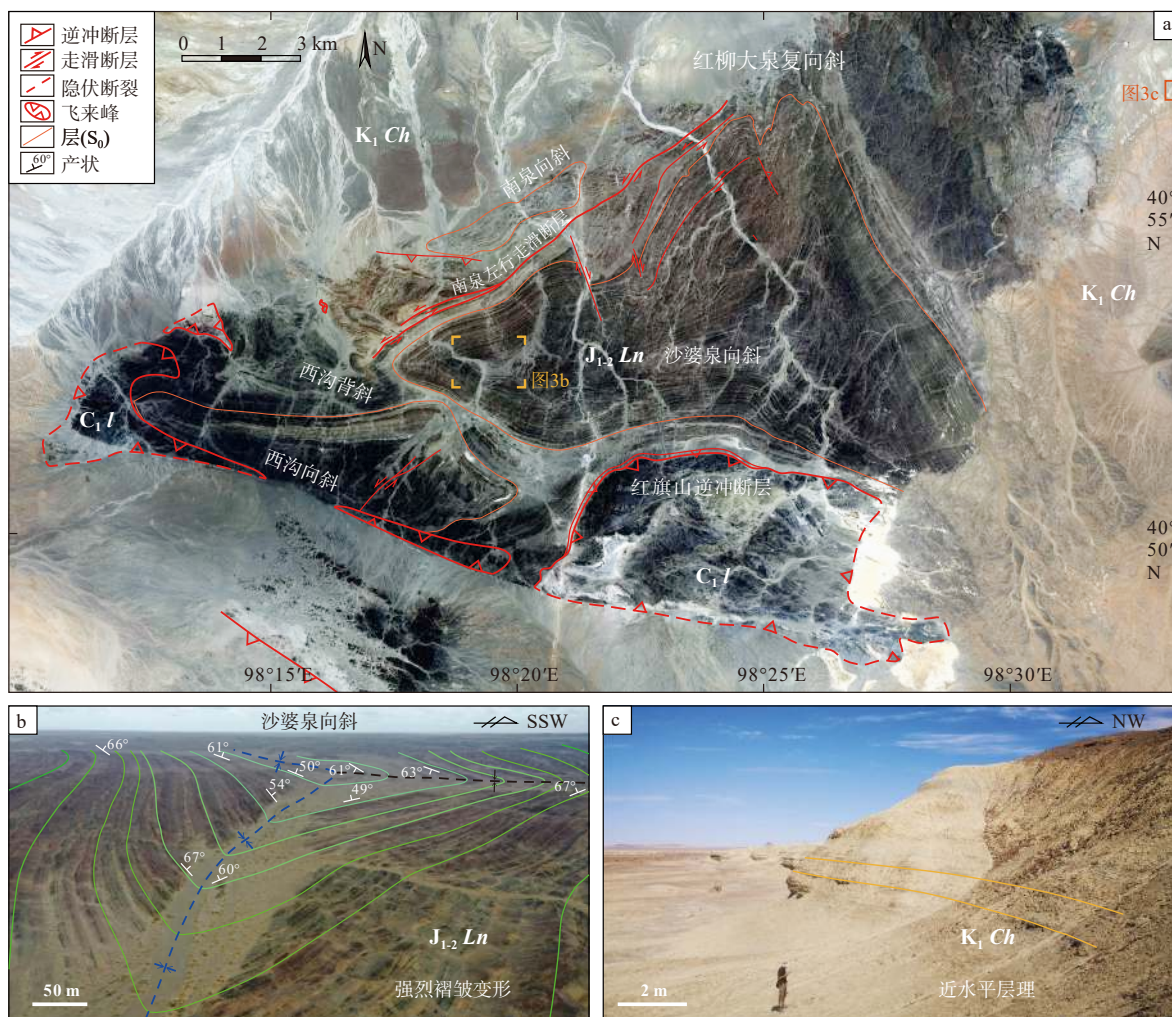
2021), 或部分逆冲断层被晚期的正断层或大型拆离断层所切割, 指示区域性的伸展构造变形 (Zheng et al., 1996; Webb et al., 1999; Graham et al., 2001; Meng, 2003; Meng et al., 2003)。北山地区新生代的构造活动很弱 (陈柏林等, 2003; Cunningham, 2013; 云龙等, 2021)。

2 构造变形分析

2.1 卫星影像解译与野外地质观察

在北山东南部红柳大泉地区, 下一中侏罗统龙凤山群 ($J_{1-2}Ln$) 广泛分布 (图 2a)。其南侧北西西—

南东东走向的红旗山逆冲断层将下石炭统柳园组及早二叠世花岗岩体往北逆冲推覆至下一中侏罗统之上 (图 2a, 图 3a)。同时, 一系列千米级别的构造盆地和短轴背斜/穹隆构造造成下一中侏罗统发生强烈褶皱变形, 如红柳大泉复向斜 (图 3a, 图 3b), 其形状呈不规则的三角形、拉长的椭圆形及变形虫等, 其枢纽沿多个方向展布, 包括北北西向、东西向、北西向和北南向 (图 2a, 图 3a)。此外, 褶皱系统内部还发育一些次级的走滑断层 (图 3a)。这些逆冲断层、褶皱及走滑断层等共同组成了红柳大泉褶皱-冲断系统 (Liu et al., 2023)。上述结果表明, 北山东南部晚中生代经历了强烈的挤压构造变形。



a—谷歌地球卫星影像的构造解译图; b—发育于下一中侏罗统内部的挤压构造变形的野外照片; c—下白垩统的野外照片

图 3 北山东南部红柳大泉地区构造解译图与野外照片 (Liu et al., 2023)

Fig. 3 Google Earth satellite image and field photos showing the structural deformation of the $J_{1-2}Ln$ and K_1Ch groups in the Hongliudaquan area, southeastern Beishan Range (Liu et al., 2023)

(a) Google Earth satellite image showing the structural deformation; (b) Field photo of the $J_{1-2}Ln$ group; (c) Field photo of the K_1Ch group

下白垩统赤金堡群(K_1Ch)大面积出露于北山东南部(图 2a, 图 3a)。相较于强烈挤压变形的下一中侏罗统(图 3b), 早白垩世地层的变形程度很弱(图 3c)。同时, 研究区地形起伏低且地表覆盖较严重, 很难通过野外地质观察揭示早白垩世总口子盆地的构造样式及属性等信息。

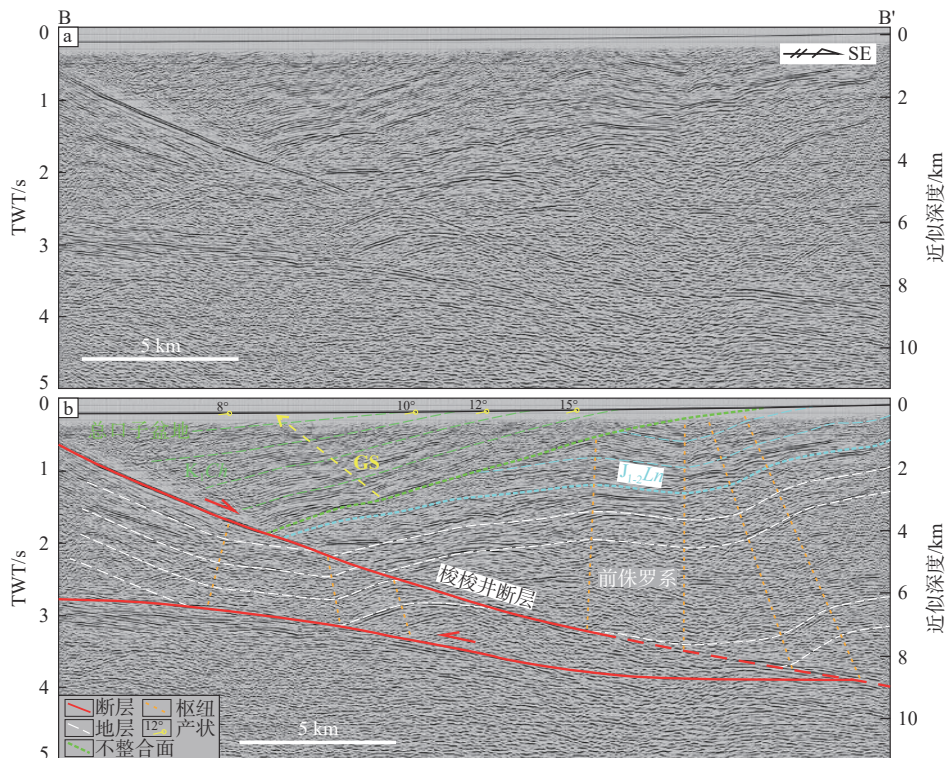
2.2 二维反射地震剖面解释

为解决上述问题, 在北山东南部红柳大泉地区选取了 3 条北西—南东走向的二维反射地震剖面。在对地震剖面进行解释时, 首先将地震剖面由时间域剖面转为深度域剖面。随后, 按照地震剖面构造解释的基本步骤开展解释工作, 具体步骤见陈宣华等(2010)。基于上述工作, 在北山东南部识别出了 2 条早白垩世伸展正断层, 即梭梭井断层与五道明断层(图 2a)。

2.2.1 正断层与断陷盆地

在地震剖面 BB' 上, 梭梭井断层的地震反射特征为一个强反射界面, 其延续性很好, 从双程走时

(TWT) ~ 0.6 s 往下延续至 ~ 3.7 s(图 4a)。在强反射界面之上, 下白垩统赤金堡群(K_1Ch)表现为弱—中等振幅、连续性较好、平行反射, 其同相轴与强反射界面呈大角度的斜交关系(图 4b)。下一中侏罗统龙凤山群($J_{1-2}Ln$)的反射特征以中等振幅、中等连续、亚平行为主。同时, 下白垩统与下一中侏罗统之间发育一个角度不整合面(图 4b)。在强反射界面之下, 前侏罗系的反射特征以中等振幅—杂乱反射为主, 其同相轴的连续性一般(图 4b)。在地震剖面 BB' 和 CC' 上, 梭梭井断层的地震反射特征类似(图 4, 图 5)。在地震剖面 DD' 上没有发育类似梭梭井断层的强反射界面(图 6a), 在剖面西段下白垩统不整合于下一中侏罗统及前侏罗系之上(图 2a, 图 6b)。该不整合面从近地表(~ 0 s)往下延伸至 ~ 1.5 s(图 6b); 剖面东段发育一个较弱的断层反射界面(图 6a), 其与下白垩统反射同相轴呈大角度的斜交关系, 该断层反射界面为五道明断层(图 6b)。



a—未解释地震剖面 BB'; b—BB'地震剖面的构造解释图

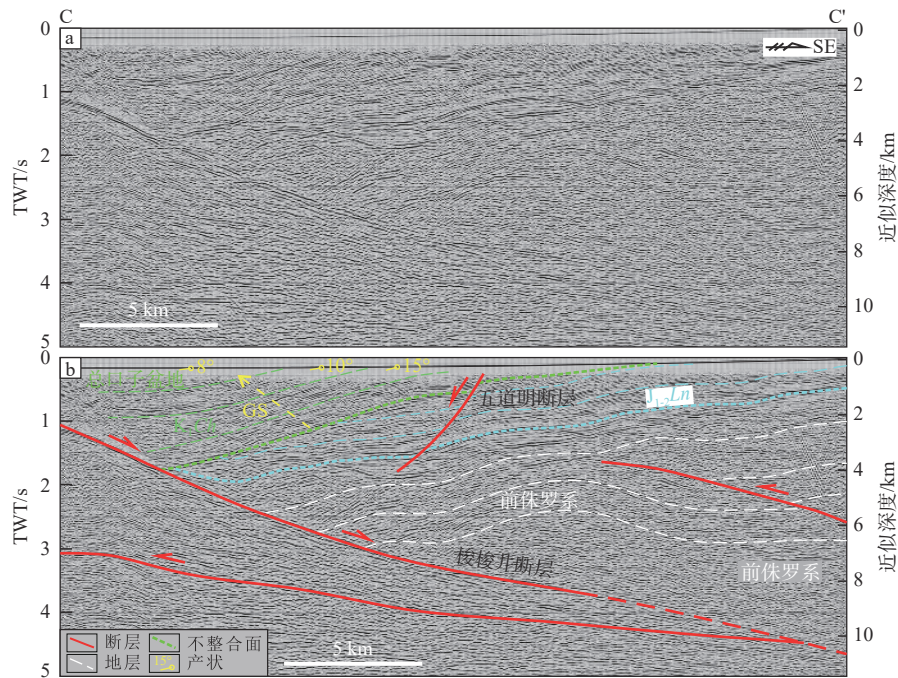
图 4 地震剖面 BB' 及其构造解释特征 (剖面位置见图 2)

Fig. 4 Seismic profile BB' and its structural interpretation (Profile location is shown in Fig. 2)

(a) Uninterpreted seismic profile BB'; (b) Structural interpretation map of seismic profile BB'

在地震剖面 BB' 和 CC' 上, 梭梭井断层倾向南东, 视倾角从剖面浅部至深部由 $30^\circ \sim 24^\circ$ 逐渐变为

$15^\circ \sim 10^\circ$, 并且其往下延伸至 $8 \sim 10$ km(图 4b, 图 5b)。梭梭井断层上盘由 K_1Ch 、 $J_{1-2}Ln$ 及前侏罗系组成, 下

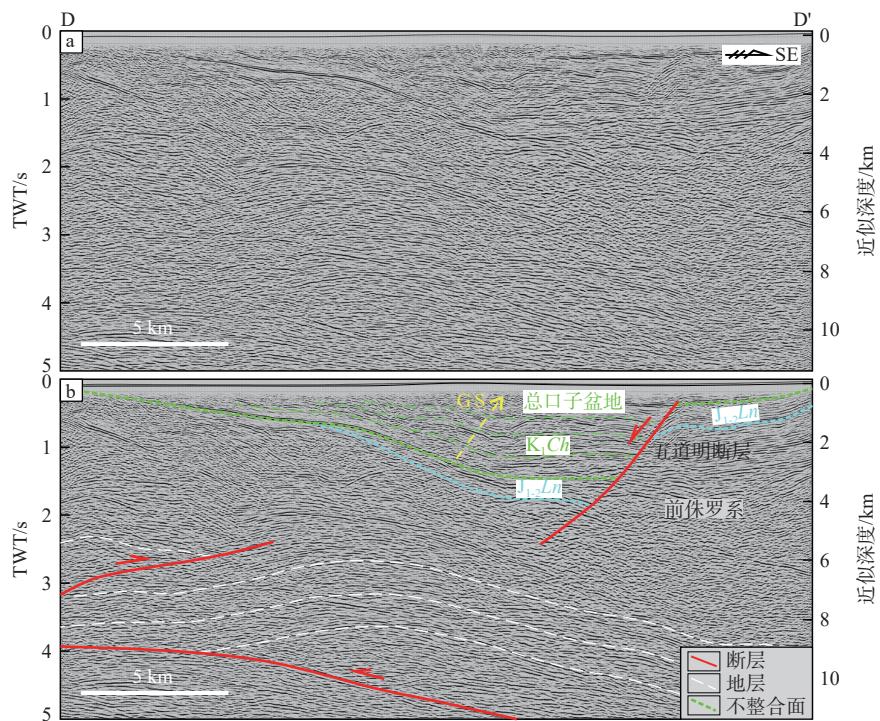


a—未解释地震剖面 CC'; b—CC'地震剖面的构造解释图

图 5 地震剖面 CC'及其构造解释特征 (剖面位置见图 2)

Fig. 5 Seismic profile CC' and its structural interpretation (Profile location is shown in Fig. 2)

(a) Uninterpreted seismic profile CC'; (b) Structural interpretation map of seismic profile CC'



a—未解释地震剖面 DD'; b—DD'地震剖面的构造解释图

图 6 地震剖面 DD'及其构造解释特征 (剖面位置见图 2)

Fig. 6 Seismic profile DD' and its structural interpretation (Profile location is shown in Fig. 2)

(a) Uninterpreted seismic profile DD'; (b) Structural interpretation map of seismic profile DD'

盘以前侏罗纪地层为主(图 4b, 图 5b)。在地震剖面 CC'和 DD'上, 五道明断层倾向北西, 倾角明显变陡, 其往下延伸至 5~6 km。五道明断层上盘以下白垩统为主, 下盘为下一中侏罗统与前侏罗系(图 6b)。因此, 梭梭井断层为一条低角度的铲式正断层(图 4b), 五道明断层为一条高角度正断层(图 6b)。平面上, 梭梭井断层的走向为北东—北东东向, 其长度超过 30 km(图 2a)。五道明断层的走向为北东向, 其长度为 10~15 km(图 2a)。

梭梭井断层和五道明断层分别限定了早白垩世总口子盆地的北西和南东边界(图 2a), 二者共同控制了早白垩世总口子盆地的形成和演化。总口子盆地在地震剖面 BB'和 DD'上表现为“半地堑”样式(图 4b, 图 6b), 而在 CC'剖面表现为“地堑”样式(图 5b)。该断陷盆地内 K_1Ch 最大沉积厚度为 3~4 km(图 4b, 图 6b)。

2.2.2 生长地层

在地震剖面 BB'和 CC'上, 下白垩统赤金堡群(K_1Ch)的地震反射同相轴具有从深部至浅部逐渐变缓的特征, 形成了典型的沉积三角楔(图 4b, 图 5b)。梭梭井断层上盘的下白垩统地层倾角从底部往上逐渐由~15°变为~12°, 并继续往上变为~8°(图 2a, 图 4b; 甘肃省地质局和第二区域地质调查队, 1971)。五道明断层上盘 K_1Ch 的地震反射同相轴也呈三角楔形状, 表现为在靠近正断层位置 K_1Ch 的厚度最大, 而远离正断层的沉积地层厚度逐渐减薄(图 6b)。上述地震反射特征表明总口子盆地内下白垩统赤金堡群(K_1Ch)发育生长地层, 并且其形成受到梭梭井断层和五道明断层持续正断层活动的控制。因此, 梭梭井断层和五道明断层均为同沉积正断层, K_1Ch 的沉积时代即为正断层的活动时间。

通过分析北山东南部俞井子盆地和公婆泉盆地内的恐龙化石特征, Li et al.(2007)和张茜楠等(2015)将 K_1Ch 的沉积时代定为早白垩世 Aptian-Albian 期(122~100 Ma)。根据孢粉组合特征, 任文秀等(2022)将总口子盆地和公婆泉盆地 K_1Ch 的沉积时代限定在早白垩世 Hauterivian-Aptian 期(132.6~113 Ma)。此外, 在该套地层中还发现了大量双壳类、腹足类、介形类及植物类等生物化石, 其时代均为早白垩世晚期(即 Hauterivian 期以来; 杨兵等, 2020)。因此, 北山东南部下白垩统赤金堡群(K_1Ch)的沉积时代很可能为早白垩世晚期。

2.2.3 挤压构造变形

在地震剖面 BB'上, 梭梭井断层下盘发育一条早期形成的逆冲断层。该逆冲断层从 9~10 km 往上延伸至 <6 km, 并且具有断坪—断坡—断坪结构(图 4b); 该逆冲断层上盘还发育一组不对称的背斜—向斜—背斜组合。在梭梭井断层上盘 $J_{1-2}Ln$ 已经发生褶皱变形, 形成一组背斜—向斜—背斜组合(图 4b)。在伸展变形之前, 这 2 套背斜—向斜—背斜组合很可能属于同一个褶皱系统, 在后期伸展变形过程中, 梭梭井断层将其错断。在地震剖面 CC'与 DD'上, 梭梭井断层与五道明断层也均切割了早期形成的挤压构造变形(图 5b, 图 6b)。

3 磷灰石裂变径迹测年

3.1 样品与实验方法

为更好地揭示北山东南部晚中生代断裂系统的发育过程和构造—热演化历史, 在梭梭井断层下盘采集了 3 块岩石样品, 用于开展磷灰石裂变径迹测试。其中, 样品 HLK-1 采集自志留纪花岗岩体, 采样位置靠近梭梭井断层; 其余 2 块样品(HLK-2、HLK-4)分别为 $J_{1-2}Ln$ 花岗岩砾石和含砾粗砂岩, 采样位置逐渐远离梭梭井断层(图 2a)。

3 块样品的磷灰石裂变径迹测试工作在中国地震局地质研究所实验测试中心裂变径迹实验室完成。实验首先将筛选、分离、提纯的磷灰石颗粒用环氧树脂固定在光片上, 进行研磨和抛光处理, 使矿物颗粒内表面露出。随后, 在 21℃ 恒温条件下, 用 5.5% 的 HNO_3 蚀刻磷灰石 20 s, 揭示出自发裂变径迹。实验采用外探测器法确定磷灰石裂变径迹年龄。将光片紧贴在低 U 白云母外探测片上, 并连同美国国家标准 CN5 标准玻璃一起放入美国俄勒冈 ^{235}U 中子活化反应堆进行辐照。照射完成后, 将白云母外探测器片在室温条件下用 40%HF 蚀刻 40 min 后, 揭示出白云母诱发径迹。径迹统计使用 Zeiss Axioplan 2 显微镜, 在放大 1000 倍干镜条件下完成。年龄计算采用 IUGS 推荐的 Zeta 标定方法(Hurford and Green, 1983), 所用 Zeta 常数为 353.0 ± 20 。年龄校正采用 Durango 磷灰石(31.4 ± 0.5 Ma)作为标准。

3.2 测试结果

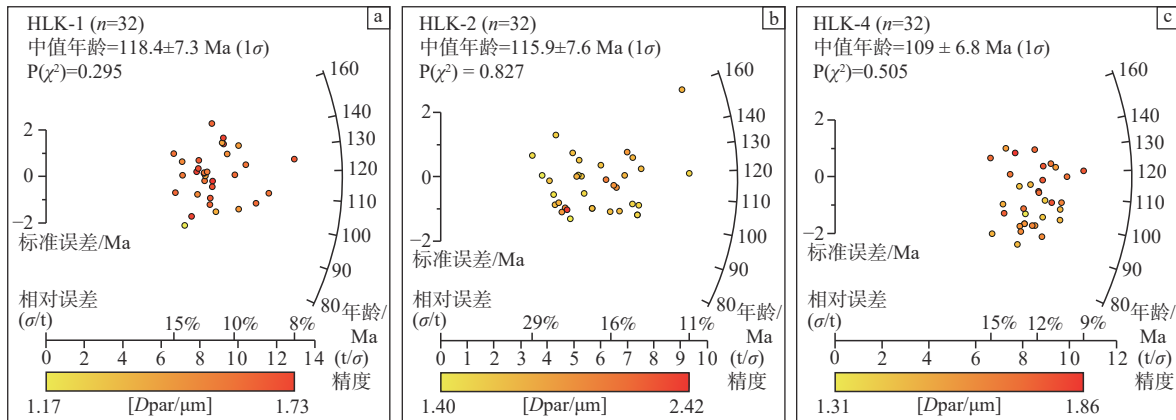
测试结果显示, 3 块样品的磷灰石裂变径迹中值年龄范围为 $118.4 \pm 7.3 \sim 109 \pm 6.8$ Ma(表 1, 图 7), 均处在早白垩世晚期, 并且均小于采样层位的岩体侵

表1 北山东南部红柳大泉地区磷灰石裂变径迹测试结果

Table 1 Apatite fission track results in the Hongliudaquan area, southeastern Beishan Range

样品编号	采样位置/ 高程	N_c	$\rho_d \times 10^6$ $\text{cm}^{-2}(N_d)$	$\rho_s \times 10^6$ $\text{cm}^{-2}(N_s)$	$\rho_u \times 10^6$ $\text{cm}^{-2}(N_u)$	U/ $\times 10^6$	$P(\chi^2)$ / %	平均径迹长度/ ($\mu\text{m} \pm 1\text{SD}$)(N_f /条)	平均 Dpar/ μm	年龄/Ma
HLK-1	40°56'20"N, 98°31'11"E/1473 m	32	1.173 (4274)	1.6434 (3983)	2.8465 (6899)	29.14	29.5	12.72±1.39 (127)	1.5	118.4±7.3
HLK-2	40°55'17"N, 98°30'28"E/1481 m	32	1.186 (4321)	1.4263 (1765)	2.5536 (3160)	25.68	82.7	12.96±0.25 (105)	1.6	115.9±7.6
HLK-4	40°53'34"N, 98°28'41"E/1517 m	32	1.18 (4297)	2.4206 (3555)	4.4873 (6737)	45.16	50.5	13.01±1.46 (111)	1.57	109±6.8

注: N_c —样品测试的磷灰石颗粒数; ρ_d —用SRM612计量剂计算的白云母外探测器的诱发径迹密度; N_d —诱发径迹总数; ρ_s —自发径迹密度; N_s —自发径迹总数; ρ_u —用晶体分析计算的白云母外探测器的诱发径迹密度; N_u —诱发径迹总数; $P(\chi^2)$ —检验所有单颗粒年龄正态分布置信度的最值(Galbraith, 1981); Dpar—与结晶c轴平行、与抛光面相交的裂变径迹蚀刻象的最大直径; N_f —统计的封闭径迹的条数



a—HLK-1 样品磷灰石裂变径迹年龄雷达图; b—HLK-2 样品磷灰石裂变径迹年龄雷达图; c—HLK-4 样品磷灰石裂变径迹年龄雷达图

$P(\chi^2)$ —检验所有单颗粒年龄正态分布置信度的最值; Dpar—与结晶C轴平行、与抛光面相交的裂变径迹蚀刻象的最大直径; $n=32$ —样品测试的磷灰石颗粒数

图7 北山东南部磷灰石裂变径迹年龄的雷达分布图(采用 RadialPlotter 软件绘制; Vermeesch, 2009)

Fig. 7 Radial plots of apatite fission track ages (plotted from RadialPlotter by Vermeesch, 2009) in the southeastern Beishan Range

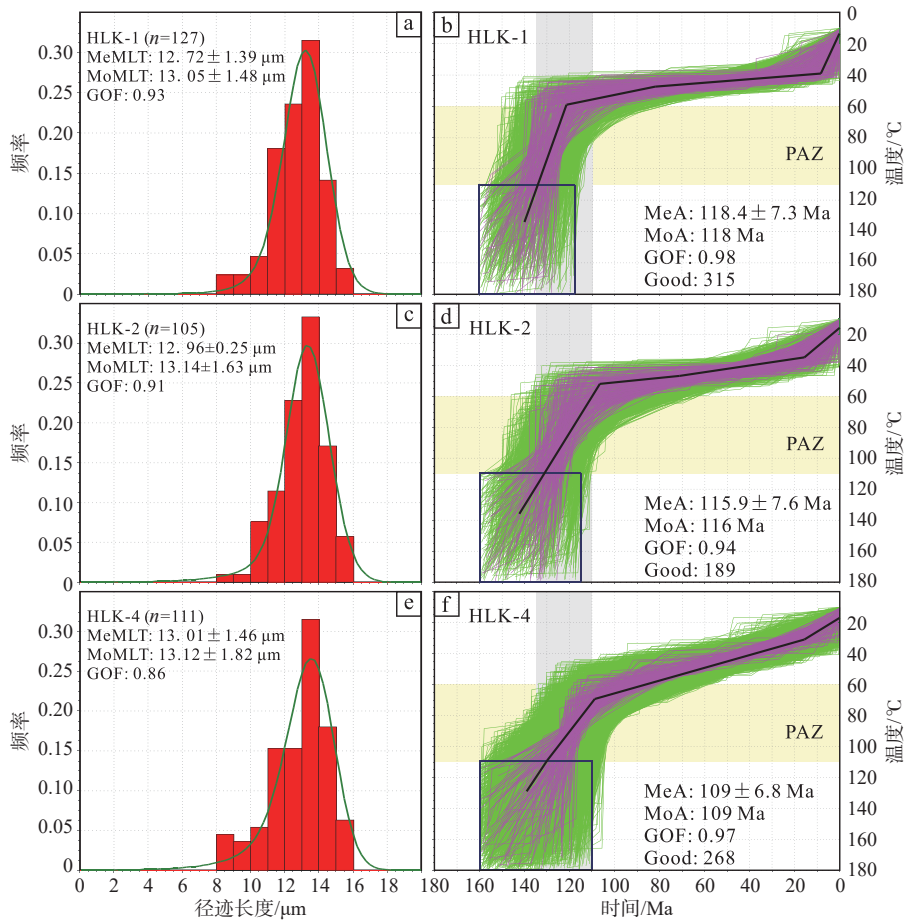
位年龄或地层沉积年龄。因此, 这些样品的磷灰石裂变径迹年龄记录了后期构造抬升和剥露过程中的冷却年龄。其中, HLK-1 样品的磷灰石裂变径迹年龄较分散($P(\chi^2)$ (检验所有单颗粒年龄正态分布置信度的最值)为 0.295, 中值年龄为 118.4±7.3 Ma(图 7a); 围限平均径迹长度为 12.72±1.39 μm (图 8a), 平均 Dpar 值为 1.5 μm 。样品 HLK-2 的磷灰石裂变径迹年龄范围为 137~95 Ma($P(\chi^2)=0.827$), 中值年龄为 115.9±7.6 Ma(图 7b); 围限平均径迹长度为 12.90±0.25 μm (图 8c), 平均 Dpar 值为 1.6 μm 。样品 HLK-4 的磷灰石裂变径迹年龄范围为 131~96 Ma($P(\chi^2)=0.505$), 中值年龄为 109±6.8 Ma(图 7c); 围限平均径迹长度为 13.01±1.46 μm (图 8e), 平均 Dpar 值为 1.57 μm 。由此可见, 随着与梭梭井断裂距离的增加, 3 块样品的磷灰石裂变径迹中值年龄逐渐变小; 另外, 较长的围限平均径迹长度值(12.72~13.01 μm)可能反映出样品以较高的冷却速率穿过部分退火带, 相似的平均 Dpar 值(1.5~1.6 μm)可能与 3 块样

品经历了相同的退火行为有关。

3.3 热史模拟

利用 Hefty 软件(v.1.9.1)和 Ketcham et al.(2007)多元退火动力学模型及蒙特卡洛逼近法, 对磷灰石裂变径迹开展了热史模拟。模拟的基础数据为实测数据, 模拟的约束条件: 初始温度设定在 180~120 $^{\circ}\text{C}$, 这一温度可使磷灰石裂变径迹完全退火(Green et al., 1989); 实测的磷灰石裂变径迹年龄区间, 使得样品处于部分退火带(110~60 $^{\circ}\text{C}$; Ketcham et al., 2007, 2009); 现今状态为地表温度(20±5 $^{\circ}\text{C}$)。每个样品反演模拟 10000 次, 这将允许计算模型统计数据并获得具有代表性的“预期”热史模型。

模拟结果显示, 梭梭井断层下盘 3 块样品经历了相似的冷却过程(图 8)。在 ~132 Ma 之前, 所有样品均处于完全退火状态(受热温度高于 120 $^{\circ}\text{C}$), 随后发生快速冷却。其中, 样品 HLK-1 于 ~132 Ma 冷却至磷灰石裂变径迹封闭温度之下(~110±10



PAZ—部分退火带; MeA—实验测试年龄; MoA—模拟年龄; GOF—模拟与实测的拟合度; Good—模拟结果为“好”的热史路径条数; MeMLT—实验测试平均径迹长度; MoMLT—模拟平均径迹长度; —最佳热史路径 —好的热史路径 —可接受的热史路径
 黑色实线代表最佳热史路径, 紫色实线代表“好”的热史路径 (GOF>0.8), 绿色实线代表“可接受”的热史路径 (0.8>GOF>0.05)
 a—HLK-1样品磷灰石裂变径迹长度直方分布图; d—HLK-1样品磷灰石裂变径迹热史模拟图;
 b—HLK-2样品磷灰石裂变径迹长度直方分布图; e—HLK-2样品磷灰石裂变径迹热史模拟图;
 c—HLK-4样品磷灰石裂变径迹长度直方分布图; f—HLK-4样品磷灰石裂变径迹热史模拟图

图8 北山东南部磷灰石裂变径迹长度分布图、热史模拟结果及径迹长度实测与模拟结果对比图(热史模拟采用HeFTy软件)

Fig. 8 Length distribution of apatite fission tracks, thermal history simulation results, and comparison of measured and simulated track lengths in the southeastern Beishan Range (thermal history simulation results using HeFTy software)

Solid black lines represent the optimal thermal history paths, solid purple lines represent "good" thermal history paths (GOF>0.8), and solid green lines represent "acceptable" thermal history paths (0.8>GOF>0.05).

℃), 随后快速穿过部分退火带(110~60℃), 并于~120 Ma 冷却至部分退火带之下(图8b); 在132~120 Ma 期间, 样品HLK-1的冷却温度至少为~50℃, 冷却速率超过了~4.1℃/Ma。样品HLK-2于132~130 Ma 冷却至磷灰石裂变径迹封闭温度之下, 并于~110 Ma 冷却至部分退火带之下(图8d); 期间样品HLK-2的冷却温度为~50℃, 冷却速率为~2.4℃/Ma。样品HLK-4于132~130 Ma 冷却至磷灰石裂变径迹封闭温度之下, 并于132~110 Ma 经历了快速冷却(图8f); 在~110 Ma 之后, 该样品的冷却速率明显变慢, 并于~90 Ma 冷却至部分退火带之下; 在快速冷却期间(132~110 Ma), 样品HLK-

4的冷却温度为~40℃, 冷却速率为~1.8℃/Ma。因此, 梭梭井断层下盘于早白垩世晚期(132~110 Ma)经历了快速冷却和剥露事件。

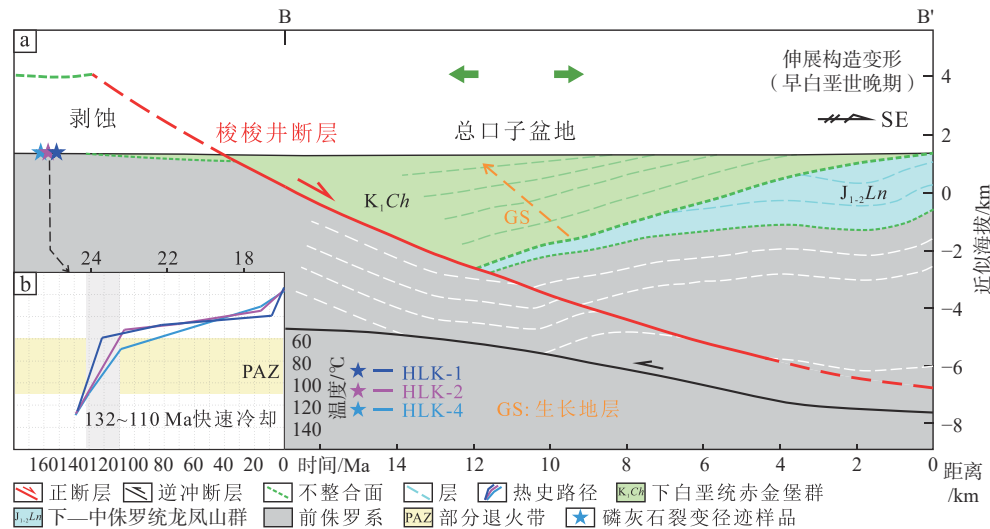
4 讨论

4.1 伸展变形时间

北山东南部梭梭井断层和五道明断层上盘的下白垩统生长地层发育, 其沉积时代为正断层的活动时间。已有研究表明, 北山东南部K₁Ch的沉积时代很可能为早白垩世Hauterivian-Albian期(132.6~100 Ma; Li et al., 2007; 张茜楠等, 2015; 张金龙等,

2017; 杨兵等, 2020; 任文秀等, 2022)。磷灰石裂变径迹热史模拟结果显示, 梭梭井断层下盘于132~110 Ma发生了快速冷却和剥露事件。这一快速冷却事件的时间与红旗山逆冲断层下盘的磷灰石(U-Th)/He冷却年龄(124~115 Ma)基本一致, 但小于红旗山逆冲断层上盘的磷灰石(U-Th)/He冷却

年龄(140~133 Ma; 图2a; Liu et al., 2023)。对比分析认为, 梭梭井断层的持续正断层活动控制了132~110 Ma下盘的剥露与快速冷却(图9)。梭梭井断层下盘快速冷却的时间与总口子盆地内 K_1Ch 的沉积时代基本一致(图9)。因此, 梭梭井断层和五道明断层的活动时间很可能为早白垩世晚期。



a—北山东南部早白垩世晚期伸展构造变形模式图; b—磷灰石裂变径迹热史模拟图

图9 北山东南部早白垩世构造-热演化模式图

Fig. 9 The Early Cretaceous tectonic-thermal evolution of the southeastern Beishan Range

(a) Late Early Cretaceous extensional structural deformation pattern in the southeastern Beishan Range; (b) Simulation map of apatite fission track thermal history

区域上, 银额盆地与东戈壁盆地内发育一系列地堑/半地堑盆地 (Graham et al., 2001; Johnson et al., 2001; Meng et al., 2003)。这些断陷盆地底部的下白垩统含132~126 Ma安山质玄武岩层 (Graham et al., 2001; 陈志鹏等, 2019)。在阿拉善东北缘, 亚干变质核杂岩主拆离断层带内的糜棱岩具129~126 Ma黑云母 $^{40}Ar/^{39}Ar$ 冷却年龄, 指示其形成时间 (Webb et al., 1999)。此外, 阿拉善至青藏高原东北缘也普遍存在126~100 Ma陆内玄武岩活动 (Zhu et al., 2008; Li et al., 2013; Hui et al., 2021), 地球化学特征指示其为陆内伸展与岩石圈减薄的产物 (Hui et al., 2021)。上述正断层、断陷盆地、变质核杂岩及玄武岩均形成于早白垩世区域伸展作用过程中。因此, 北山东南部及周缘地区早白垩世伸展构造变形很可能开始于133~129 Ma, 并持续到了早白垩世末期。

4.2 挤压-伸展构造转换及区域意义

中侏罗世晚期—早白垩世早期, 强烈的陆内挤

压变形将古生代—早中生代北山造山带构造活化, 形成了晚中生代北山褶皱-冲断带。期间, 北山逆冲推覆构造发生了大规模(>120 km)由南往北的逆冲运动 (左国朝等, 1992; Zheng et al., 1996), 星星峡断层也经历了右行斜向逆冲运动 (Zhang and Cunningham, 2012)。在北山东南部, 红柳大泉褶皱-冲断系统沿2个方向同时发生水平缩短, 并吸收了至少~27%的北南向的地壳缩短量, 这使得北山地壳被增厚至50~60 km (Liu et al., 2023)。该期陆内挤压变形事件普遍被认为是亚洲陆缘多板块汇聚作用远程传递的结果, 包括蒙古-鄂霍茨克洋闭合、拉萨与羌塘地体碰撞及古太平洋板块向西俯冲 (Zheng et al., 1996; Dumitru and Hendrix, 2001; Guynn et al., 2006; Zhang and Cunningham, 2012; Dong et al., 2015; Van der Voo et al., 2015; Ma et al., 2017; Ma and Xu, 2021)。

早白垩世晚期 (133~129 Ma), 区域构造变形

体制发生了重大转变,北山地区及周缘开始转入北西—南东向伸展变形阶段。在北山东南部,梭梭井断层与五道明断层切割了红柳大泉褶皱—冲断系统。其中,梭梭井断层具铲式正断层的特征,往下很可能汇入了一条拆离断层中(脆—韧转换带;图10)。这2条正断层分别限定了早白垩世总口子盆地的北西和南东边界。区域上,阿拉善和南蒙古地区发育一系列同时期的地堑/半地堑盆地(图11;

Graham et al., 2001; Meng, 2003; Wang et al., 2011), 其共同组成了早白垩世的伸展盆地系统,如银额盆地(Graham et al., 2001; Meng, 2003; Meng et al., 2003; Hui et al., 2021)。其中,部分断陷盆地具有超拆离盆地的性质,与亚干变质核杂岩及大型拆离断层的形成密切相关(Webb et al., 1999; Meng et al., 2003)。至早白垩世末期,早期增厚的北山地壳(50~60 km)被不同程度地拉伸减薄。

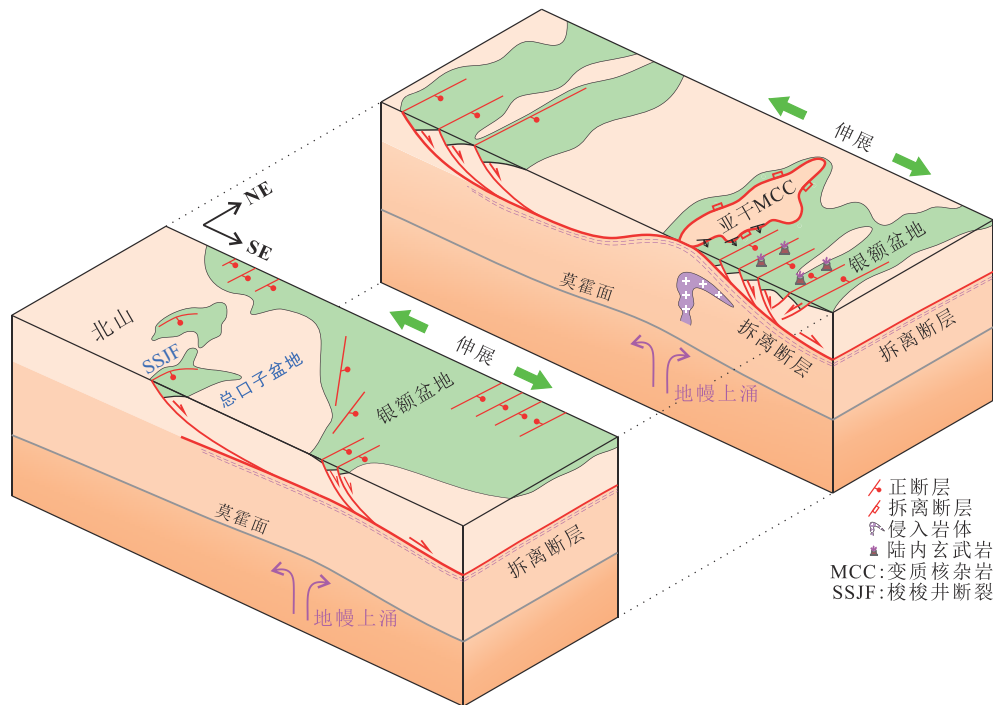


图10 北山地区及周缘早白垩世伸展构造变形模式图 (Webb et al., 1999; Meng et al., 2003; Hui et al., 2021; Zuo et al., 2020)

Fig. 10 Tectonic model of the Early Cretaceous extensional deformation in the Beishan Range and its surrounding regions (modified from Webb et al., 1999; Meng et al., 2003; Hui et al., 2021; Zuo et al., 2020)

晚中生代中亚造山带南缘普遍经历了挤压—伸展构造的转换 (Webb et al., 1999; Johnson et al., 2001; Davis et al., 2002; Cunningham et al., 2009; Wang et al., 2011, 2015; Zhang et al., 2012, 2021a, 2021b; Zhu et al., 2015; Lin and Wei, 2018)。针对这一构造转换,部分学者认为北山、南蒙古和阿拉善地区伸展构造变形开始于晚侏罗世 (~155 Ma; Zheng et al., 1996; Graham et al., 2001; Johnson et al., 2001; Meng, 2003; Meng et al., 2003)。综合已有成果,研究认为北山地区及周缘晚中生代陆内挤压—伸展构造转换很可能发生在早白垩世晚期 (133~129Ma)。这一构造转换的时间与华北克拉通北缘燕山—阴山构造带及整个东亚地区大规模伸展构造变形的启动时间 (~135

Ma)基本一致 (Davis et al., 2001; 杨进辉等, 2006; Davis and Darby, 2010; 张长厚等, 2011; 刘少峰等, 2018; Wang et al., 2018; Lin et al., 2019, 2020)。同时,伸展构造变形的方向也具有—致性(北西—南东向; Wang et al., 2011; Lin and Wei, 2018; 林伟和李金雁, 2021)。因此,晚中生代中亚造山带南缘很可能经历了统一的陆内挤压—伸展构造转换过程(图11)。

4.3 动力学模式

早白垩世,亚洲陆缘多板块汇聚体系的深部过程和构造体制发生了重大转变 (Lin et al., 2013; Dong et al., 2015; Zhu et al., 2015; 张岳桥和董树文, 2019; Ma and Xu, 2021; Chen et al., 2022)。一方面,随着蒙古—鄂霍茨克洋最终关闭,蒙古—鄂霍茨克造山带转

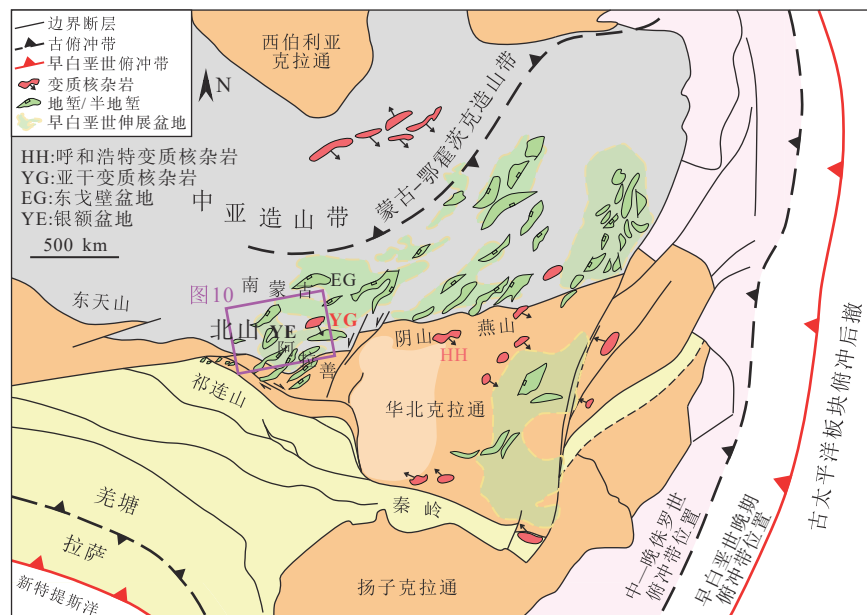


图 11 早白垩世中亚和东亚地区伸展构造变形 (包括变质核杂岩和地堑/半地堑盆地等) 与亚洲陆缘多板块汇聚体系的关系图 (Meng, 2003; Meng et al., 2003, Wang et al., 2011, 2015; 任纪瞬等, 2013; Lin and Wei, 2018; Ma and Xu, 2021)

Fig. 11 Distribution of Early Cretaceous extensional structures in the central and eastern Asian continent, including metamorphic core complexes and graben/half-graben basins (modified from Meng, 2003; Meng et al., 2003; Wang et al., 2011, 2015; Ren et al., 2013; Lin et al., 2018; Ma and Xu, 2021)

入后造山演化阶段 (Meng, 2003; Wang et al., 2011, 2015; 董树文等, 2019)。同时, 拉萨地体与羌塘地体的碰撞拼贴也已基本完成 (Fan et al., 2015, 2018; Peng et al., 2020)。因此, 南北向汇聚作用力逐渐减弱 (Meng, 2003; 董树文等, 2019; 林伟和李金雁, 2021; Chen et al., 2022)。另一方面, 俯冲的古太平洋板块开始往东回撤, 其俯冲角度也逐渐变陡 (图 11; 朱日祥等, 2012; 朱日祥, 2018; 张岳桥和董树文, 2019; Ma and Xu, 2021)。这一重大转变导致区域挤压应力减小甚至消失, 进而引发强烈增厚地壳的重力垮塌。同时, 在中亚造山带南缘曾发现众多早白垩世陆内玄武岩, 表明大量幔源物质加入到了中亚大陆地壳之中, 指示地幔上涌 (Graham et al., 2001; Zhu et al., 2008; Li et al., 2013; 陈志鹏等, 2019; Hui et al., 2021)。上涌的地幔物质持续作用于岩石圈底部, 为早白垩世区域伸展作用提供了水平拉伸力 (Raimondo et al., 2014; 张进江和黄天立, 2019; Zuo et al., 2020; Hui et al., 2021)。早白垩世古太平洋板块往南东方向发生了大规模俯冲后撤 (图 11), 引发了东亚地区软流圈地幔侧向流动或非稳态地幔流 (朱日祥等, 2011, 2012; 朱日祥, 2018; Ma and Xu, 2021)。

这一深部地球动力学过程很可能打破了中亚地区地幔对流的原有平衡, 并诱发局部软流圈地幔上涌 (Hui et al., 2021; Zuo et al., 2020)。因此, 亚洲陆缘多板块汇聚体系的重大转变诱发了增厚地壳的重力垮塌及局部地幔上涌, 进而导致了中亚造山带南缘早白垩世的区域伸展作用 (图 10, 图 11)。

5 结论

(1) 在北山东南部, 一系列逆冲断层与褶皱构造造成下一中侏罗统发生强烈挤压变形。二维反射地震剖面揭示出 2 条早白垩世伸展正断层, 其中梭梭井断层为倾向南东的低角度铲式正断层, 五道明断层为倾向北西的高角度正断层。二者均切割了中侏罗世晚期—早白垩世早期的褶皱—冲断系统, 指示挤压—伸展构造的转换。梭梭井断层与五道明断层分别限定了早白垩世总口子盆地的北西和南东边界, 使得其发育“地堑”样式。总口子盆地内下白垩统生长地层发育, 表明梭梭井断层与五道明断层的活动时间为早白垩世晚期。

(2) 磷灰石裂变径迹中值年龄范围为 118.4 ± 7.3

Ma~109±6.8 Ma。热史模拟结果显示, 梭梭井断层下盘于 132~110 Ma 经历了快速冷却和剥露事件, 该事件与梭梭井断层的持续正断层活动密切相关。

(3) 北山东南部晚中生代陆内挤-伸展构造转换的时间很可能为 133~129 Ma(早白垩世晚期)。增厚地壳的重力垮塌与局部地幔上涌共同导致了中亚造山带南缘早白垩世北西-南东向伸展构造变形。

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