Postglacial sea-level rise in south Portugal as recorded in Guadiana Estuary

T. Boski¹, D. Moura¹, S. Camacho¹, D. Duarte¹, D. B. Scott², C. Veiga-Pires¹, P. Pedro¹, P. Santana¹

¹UCTRA-CIMA, Universidade do Algarve, Campus de Gambelas, 8000 FARO, PORTUGAL.

e-mail (Boski): tboski@ualg.pt

e-mail (Moura): dmoura@ualg.pt

e-mail (Duarte):dduarte@ualg.pt

e-mail (Veiga-Pires): cvpires@ualg.pt e-mail (Pedro):ppedro@ualg.pt

e-mail (Santana):psantana@ualg.pt

² Centre of Marine Geology, Dalhousie University, Halifax. Nova Scotia, CANADA e-mail: dbscott@is.dal.ca

The Guadiana River Estuary is located in the terminal part of a deeply incised river valley, which accumulated several tens of meters of sediments during the Holocenic transgression. Five cored boreholes (see Fig.1 for localisation) that reached the pre-Holocenic substratum were drilled recently in order to recognize the architecture of sedimentary facies and to quantify the accumulation of organic carbon trapped in sediments during the valley infilling by marine waters. It was assumed that due to structural constraints imposed by Palaeozoic and Mezozoic substratum, the main estuarine channel did not change its position significantly. Consequently borehole locations were chosen in order to represent different sedimentary environments in the estuary: proximity to the main channel (CM1 and CM3), external sea facing (CM4) and lagoonal (CM2 and CM5) environments. In all five boreholes the Holocene was found to overlay Pleistocene fluvial gravels at depth of 39.2m in CM1, 10m in CM2, 31m in CM3, 19.2m in CM4 and 50.8m in CM5. These depths reflect the pre-Holocenic topography of Guadiana Valley: the basal gravels correspond respectively to Pleistocene terrace levels, in boreholes CM1, CM3 and CM4, to shallow Jurassic platform in CM2 and to the bottom of Beliche River (Guadiana tributary) valley in CM5.

From Table 1 which resumes results of datings so far done, it appears that the entire Holocenic sedimentary history is recorded within the deeper portions of Pleistocene river valleys (CM5 and CM1 boreholes), where the initial most rapid part of transgression, resulted in the deposition of monotonous clay sequence reported in the works of Dabrio et al. (1995) and Goy et al. (1996) from Tinto–Odiel Estuary in Spain. This type of sedimentation, i.e. from the upper intertidal regime, prevailed until present day in the confined areas of CM2 and CM5 boreholes. In the areas of CM1 and CM3 the basal clay portion is followed by the succession of meander bar sequences and on top by a lateral bar sequence.

In the area of CM4 which is the most sea exposed, the basal clay sequence is thin or almost lacking (yet to be checked by datings), due to the elevation of Pleistocene surface. The coarse/medium sand sequence represents depositional environment of coastal bars and in the top 3m, dunes.

The microfaunistic analyses of CM1 and CM3 yielded to the identification of two species associations, with a very distinct Shannon's index. This one abruptly changed from 0.03 to 2.28 at 8430 BP in CM1 and from 0.00 to 2.97 at 9470 BP in CM3. These associations testify the evolution from saltmarsh environment to a medium intertidal environment, between 8430±380 BP and 9470±250 BP.

The species association dominated by A. beccarii and H. germanica, did not undergo alteration along borehole CM2 sedimentary columm.

The major marine influence was recorded in CM4 through a high plancktonic/benthic foramineferous ratio, which begins to increase significantly around 6250±250 BP.

References

Borrego, J.; Morales, J. A. & Pendon, J. G. (1993) Holocene filling of an estuarine lagoon along the mesotidal coast of Huelva: The Piedras Tiver mouth, southwestern Spain. Journal of Coast. Res. 9, 242-254.

Goy, J. L. ; Zazo, C.; Somoza, L.; Dabrio, C.J. Lario, J.; Borja, F.; Sierro, F. J. & Flores, J. A. (1996) Global and regional factors controlling changes of coastlines in South Iberia (Spain) during Holocene. Quaternary Science Reviews, 15, 773–780.

| Borehole | Sample depth(cm) | Age BP(yr.) | δ ¹³ C(‰) | Material | Method |
|----------|------------------|-------------|----------------------|----------------|-------------|
| CM1 | 824 | 5020±310 | 1.1 | shells | radiometric |
| CM1 | 1530 | 4600±40 | -26.9 | organic matter | AMS |
| CM1 | 1712 | 6210±220 | -25.9 | peat | radiometric |
| CM1 | 1860 | 6205±40 | 2.08 | shells | AMS |
| CM1 | 2127 | 7590±100 | 20 | shells | radiometri |
| CM1 | 2850 | 8430±380 | | shells | radiometri |
| CM2 | 200 | 3080±100 | -27.3 | organic matter | radiometri |
| CM2 | 715 | 5950±190 | 0.4 | shells | radiometri |
| CM2 | 854 | 10130±200 | -5.98 | shells | radiometri |
| CM3 | 459 | 3300±160 | -0.6 | shells | radiometri |
| CM3 | 960 | 6710±120 | 1.6 | shells | radiometri |
| CM3 | 1452 | 7080±200 | 0.7 | shells | radiometri |
| CM3 | 2690 | 9470±250 | -22.9 | wood | radiometri |
| CM4 | 700 | 6200±340 | 1.6 | shells | radiometri |
| CM4 | 975 | 5640±90 | -0.31 | shells | radiometri |
| CM4 | 1575 | 6250±250 | 1.4 | shells | radiometri |
| CM4 | 2834 | 16980±100 | -28.5 | organic matter | AMS |

Table 1 Summary of sediment datings from 4 boreholes in Guadiana estuary.

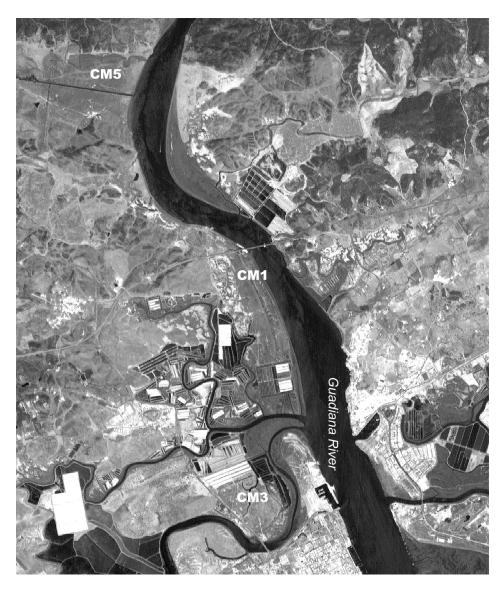


Fig. 1 Geographical localication of the five cored boreholes.