

Improved understanding of eutrophication trends, indicators and problem areas using machine learning

Deep S Banerjee¹ and Jozef Skakala¹

¹Plymouth Marine Laboratory

May 08, 2024

Improved understanding of eutrophication trends, indicators and problem areas using machine learning

Deep S. Banerjee^{1,2} and Jozef Skákala^{1,2}

¹Plymouth Marine Laboratory, PL1 3DH Plymouth, United Kingdom,

²National Centre for Earth Observation, PL1 3DH Plymouth, United Kingdom.

Correspondence: Deep S. Banerjee (dba@pml.ac.uk)

Abstract. Nitrate is an essential inorganic nutrient limiting phytoplankton growth in many marine environments. Eutrophication, often caused by nitrogen deposition, is a reoccurring problem in coastal regions, including the North-West European Shelf (NWES). Despite of their importance, nitrate observations on the NWES are difficult to obtain and thus sparse both in time and space. We demonstrate that machine learning (ML) can generate, from sparse observations, a skilled, gap-free, bi-decadal (1998-2020) surface nitrate data-set. We demonstrate that the effective resolution (scales on which the data-set is skilled) is slightly coarser than the 7 km and daily resolution of the product, but still completely sufficient to analyse nitrate dynamics on a monthly scale. With such a data-set we can address questions that would be otherwise hard to answer: (i) We show that nitrate-limited regions on the NWES, potentially vulnerable to eutrophication, extend beyond the eutrophication-problem areas already identified by the monitoring bodies (i.e. OSPAR). The newly identified regions include southern Irish coastline and parts of Irish Sea, indicating that these areas could become problematic under sub-optimal policy, or management changes. (ii) We demonstrate that bi-decadal 1998-2020 trends in coastal nitrate, responding to long-term policy-driven reduction in riverine discharge, are mostly modest with a notable exception of the Bay of Biscay. (iii) We show that winter nitrate plays relatively minor direct role in the phytoplankton bloom intensity the following spring, which can have some implications for using winter inorganic nitrogen as eutrophication indicator (as often included by OSPAR).

1 Introduction

Nitrogen is one of the most important components of organic matter, needed in relatively large concentrations, as demonstrated by the Redfield ratios (Tett et al., 1985). Despite of its large abundance (the Earth's atmosphere comprises 78% nitrogen as N₂), it is non-trivial to obtain nitrogen in forms useful for plants. As a consequence of this, nitrogen is often the most limiting nutrient for plant, or algae growth, including the coastal marine environment (Ryther and Dunstan, 1971; Board et al., 2000). Nitrogen fixation, converting atmospheric nitrogen to forms useful for life, happens through various biotic and abiotic pathways, resulting in ammonium, nitrite and nitrate (Noxon, 1976; Hill et al., 1980; Postgate, 1998; Beman et al., 2008; Voss et al., 2013). Nitrate in the ocean is the primary nutrient for phytoplankton, with phytoplankton uptake enabling nitrogen flows into higher trophic levels and various detrital and dissolved forms of organic matter. In a nitrogen-limited environment, excess nitrate concentrations, primarily originating from agricultural runoff and industrial wastewater discharge, can stimulate harmful

25 eutrophication events (Withers et al., 2014; Nazari-Sharabian et al., 2018). The thick layer of algae produced by these events
may cut oxygen ventilation at the surface and after the algae die off and sink, the decomposers may consume vast amounts
of oxygen, leading to marine hypoxia in the bottom part of the water column (Rabalais et al., 2002; Diaz and Rosenberg,
2008). Furthermore, eutrophication events are often dominated by species that produce toxins that have detrimental effects
on the marine ecosystem by causing fish kills, seafood contamination, and even posing risks to human lives (Anderson et al.,
30 2012). Additionally, high nitrate concentrations lead to the excessive production of organic matter, which, upon decomposition,
increases CO₂ concentration, contributing to ocean acidification (Doney et al., 2009). Eutrophication is a fundamental problem
in many shelf sea and coastal areas (Rabalais et al., 2009), with nitrate monitoring and predicting providing an essential tool
informing marine management and policy.

An important region, subject to eutrophication, is the North-West European Shelf (NWES). NWES is impacted by signifi-
35 cant river inputs, notably the Thames, Rhine, and Loire, which introduce substantial freshwater and nutrients into the region,
influencing salinity and water properties. Open ocean shelf exchange, especially transport of nutrients and carbon across the
shelf break, play another vital role in the NWES ecosystem dynamics (Huthnance et al., 2009). NWES has high ecological im-
portance due to its high biological productivity, underpinning significant commercial fisheries and carbon sequestration (Pauly
et al., 2002; Borges et al., 2006; Jahnke, 2010). During the 1980s, the NWES, particularly near the German Bights and the
40 Westerschelde estuary, experienced notable shifts in nutrient distribution, primarily driven by increased continental nutrient
inputs. Riverine discharges, particularly from the Rhine and Elbe, have been identified as major contributors to nutrient dy-
namics in the region (Brockmann and Eberlein, 1986; Radach, 1992), having adverse effects on the local ecosystem. However,
EU regulations set by OSPAR convention in 1992 substantially decreased the nitrate deposition into the NWES (Burson et al.,
2016).

45 The NWES nitrate concentrations are operationally simulated and predicted (Skákala et al., 2018), however, the NWES
nitrate observations are too sparse to properly constrain the simulated nitrate through data assimilation. The current operational
NWES system is mainly constrained by the much more robust satellite temperature and chlorophyll observations (Skákala et al.,
2018, 2021, 2022) and avoids assimilating nutrients entirely. Furthermore, due to its univariate nature, the operational system
fails to directly constrain most of the non-assimilated variables including nutrients. Consequently, the nitrate reanalyses and
50 forecasts produced by the operational system are known to have substantial biases, inherited from the model free run (Skákala
et al., 2018, 2022). Although the simulated physics and chlorophyll from the reanalysis validate well against observations
(Skákala et al., 2018, 2022), the nitrate NWES product is of more limited use.

In this work we develop and validate a new bi-decadal NWES nitrate product derived from the available observations using
advanced machine learning (ML) algorithms. The nitrate product is developed for the ocean surface, where nutrients have the
55 potential to most significantly drive phytoplankton growth. This is up to our knowledge the by far most complete and detailed
observation-based sea surface nitrate data-set on the NWES. Unlike the NWES operational reanalysis, the data-set validates
skillfully against the independent observations. Using our NWES nitrate product we are able to discuss several important
questions, like the impact of winter nitrate pre-conditioning on the inter-annual phytoplankton variability, identify the NWES
geographic areas limited by nitrate, or analyse trends in nitrate concentrations on the NWES. To do so, we maximise our

60 reliance on the observational data and use ML and modelling to effectively fill the large data-gaps, either through statistics, or dynamical consistency imposed by deterministic modelling.

2 Methodology

2.1 The ML model

We used as the ML model a Feed-forward Neural Network (NN) designed through the Autokeras library using a Structured
65 Data Regressor (Jin et al., 2019). This approach streamlined the process of hyperparameter optimization and model architecture discovery through an automated procedure, significantly reducing the need for manual intervention. The routine follows an iterative trial-and-error approach by looping through several possible combinations of various hyperparameters, such as learning rate, optimizers, and the number of dense layers with different combinations of nodes. It then selects the best model architecture with the highest skill score against the validation dataset and saves it for final prediction against the test data.

70 The final model architecture comprised several layers (see Fig.S1 of Supporting Information, SI): (i) the input layer with 25 nodes corresponds to the input features or predictors, (ii) a multi-category encoding layer, encoding categorical features (i.e., month and day of the year) into a numeric form that can be understood by the network, (iii) a normalization layer, which normalizes the input data to improve model training by facilitating improved model convergence time and better performance, avoiding dominance of features with larger magnitudes, and providing better stability, (iv) two dense layers with 128 and
75 256 nodes, respectively, with each of these layers being followed by a Rectified Linear Unit (ReLU) activation function to introduce non-linearity into the model, (v) dropout layer applied to prevent overfitting by randomly dropping out a fraction of neurons during the training phase, and (vi) regression head with a single node that produces the final prediction, i.e., nutrient concentration. During the training phase, the model's performance was evaluated using standard metrics, i.e., Mean Squared Error (MSE), and then estimating relative error with respect to the validation dataset.

80 2.2 Data

2.2.1 The input features

To avoid biases towards operational models, the NN model input features were always selected to be either observational data, or reanalyses of variables closely constrained by the observations. One of the main challenges in building ML for environmental applications is to combine (often sparse) data with typically inconsistent domains of coverage across variety of spatial
85 and temporal scales. The most robust observational marine data-sets are obtained through satellite optical measurements, for physics these are data such as sea surface temperature (SST), or altimetry, for biogeochemistry the most typical derived data-set is surface chlorophyll *a* concentration obtained from the ocean color (OC). Recently, new remote-sensing algorithms were developed to partition the total chlorophyll concentration into phytoplankton functional types (PFTs) based largely on phytoplankton size-classes (Brewin et al., 2010, 2017), and PFT chlorophyll products are now operationally assimilated into the
90 NWES model (Skákala et al., 2018). The reanalyses from the NWES operational model have been found to have a very close

match-ups with the assimilated observations (Skákala et al., 2018, 2022), and act as natural extensions of the observations, forming a complete data-set on a gridded domain. We have extracted a range of NN model features from a NWES physical-biogeochemistry reanalysis product (downloadable from the EU Copernicus portal, <https://marine.copernicus.eu/>, see also Kay et al. (2016)), covering the 1998-2020 period with daily and 7 km spatial resolution. The product is based on assimilating satellite SST, temperature and salinity profiles, as well as OC PFT chlorophyll, into the operational Nucleus for Ocean Modelling (NEMO, Madec et al. (2017)) model, coupled through Framework for Aquatic Biogeochemical Models (FABM, Bruggeman and Bolding (2014)) to the biogeochemistry, European Regional Seas Ecosystem Model (ERSEM, Butenschön et al. (2016)). The extracted features were for (i) SST, (ii) chlorophyll ocean surface concentrations from four PFTs, which were assimilated into the model (diatoms, microphytoplankton, nanophytoplankton and picophytoplankton), as well as for (iii) total surface phytoplankton carbon, (iv) total surface chlorophyll and (v) total surface net primary production. Although these features were selected from the reanalysis, they either correspond to the assimilated variables (SST, PFT chlorophyll), or to variables which are dynamically very close to the assimilated PFT chlorophyll (phytoplankton carbon, net primary production) and therefore well constrained by the assimilation. Additionally, we also included SST observations from the Global Ocean OSTIA product (Good et al., 2020; Donlon et al., 2012) in the input feature dataset. Interestingly our tests (not shown here) indicated that the NN model did not perform as well without this additional SST data-set, so both sources of SST information (reanalysis and OSTIA) were used.

We have also used input features describing riverine discharge into the ocean. These included riverine discharge data for oxygen and nutrient loads (i.e., nitrate, phosphate, silicate, ammonia, and oxygen) at all relevant river mouths in the NWES domain. Daily time series of river discharge are used for 1998-2017. From 2018 only climatologies were available and were used for the remaining 2018-2020 period. The daily riverine data were obtained from an updated version of the river dataset from Lenhart et al. (2010), climatology of daily discharge data were taken from the Global River Discharge Data Base and the Centre for Ecology and Hydrology (Young and Holt, 2007). Unlike a hydrodynamic model, a NN does not follow any advection mechanism or transport to carry the effect of river discharge over space and time. To account for advection the NN model would have to ideally include time-lagged riverine inputs, where the time-lag would increase with the spatial distance from the river mouth. This would hugely increase the complexity of the NN model. In this work we have decided to opt for less complex models and avoid using time-lagged NN inputs. To consider nearly-instantaneous riverine effect in such simplified scheme, we have distributed the riverine discharge data around the river discharge point sources. This was done by spatially extrapolating all the runoff variables at all the daily discharge points over a 50 km circular radius surrounding the main discharge point, by making the inputs decay inversely from the maximum at the center to a zero value at the edge (see Fig.1).

Another data-set that was used as input features into the NN model was the ERA-5 atmospheric reanalysis (Hersbach et al., 2020), which has a horizontal resolution of 0.25° . The input feature dataset includes variables such as downwelling shortwave radiation at the ocean surface, specific humidity, temperature at 2m above the ocean surface, sea level pressure, total precipitation, and zonal and meridional wind components at 10m above the ocean surface. These near-surface atmospheric drivers play a crucial role in governing and redistributing surface nitrate through air-sea interactions and atmospheric deposition

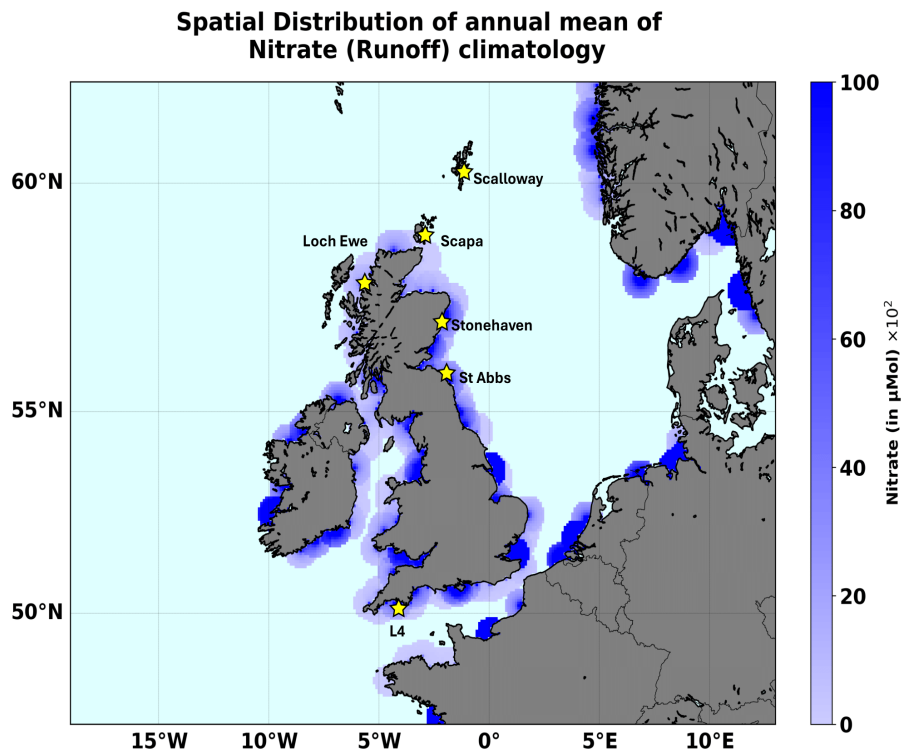


Figure 1. The areas with river input on the NWES domain. We mark the location of the stations providing the test data for this study: L4 and the five Scottish coastal stations.

125 of nitrogen, which accounts for one-third of the non-recycled nitrogen supply in the ocean and up to around 3% of annual new biological production (Duce et al., 2008).

Finally, we used as input features also structural temporal and geographic data, i.e. the time of the year (month, day), latitude, longitude, depth and bathymetry. This type of information enables the model to learn the geographic patterns in nitrate, including its seasonal climatology, substantially enhancing its predictive accuracy.

130 All the input features were considered at the same times than the predicted nitrate. The input features importance was ranked in the SHAP analysis, presented in Fig.S1 of the Supporting Information (SI). The SHAP analysis indicates that the structural input features are among the most important, followed by some atmospheric and oceanic physics features (incoming short-wave radiation, SST), which can partly account for seasonal climatology as well. Specific riverine discharge input features (e.g. of nitrate itself) are highly important too (Fig.S1 of SI), with range of biogeochemical variables (total surface chlorophyll, 135 total surface net primary production and total surface carbon) being around the middle of the importance ranking. As already

mentioned, we have tested also versions of the NN model with reduced number of input features (e.g. removing the less important features from the SHAP analysis), but the performance of such model turned out to be slightly worse.

2.2.2 The predicted nitrate data, the training and the validation process

The 1998-2018 nitrate observations were obtained from the International Council for the Exploration of the Sea (ICES) Data portal (<https://www.ices.dk>). The extracted dataset spans a geographical range from 19W to 10E in longitude and from 47N to 62N in latitude, ensuring wide representation of the dynamics and variability of the NWES region. The ICES data were obtained from a wide range of in situ measurements, e.g. by cruises, floats, moorings, or buoys. In the training and validation process, the ML model inputs were interpolated into the ICES data locations, and then the ICES data from the 1998-2015 period, containing 43572 relevant data-points, were used for training and validation of the NN model (with 80% data used for training and 20% used for validation). Finally, the 2016-2018 ICES data, containing 2984 data-points, were used as test data. The spatial coverage of all the ICES training, validation and test data is shown in Fig.2.

Several other observations were used as test data, to demonstrate the ML model skill: (i) nitrate data from the L4 station, which is operated by the Western Channel Observatory (<https://www.westernchannelobservatory.org.uk/>) and is located in the western English Channel approximately 13 km from the Plymouth Sound, providing one of the longest continuous ecological time-series in the world (Harris, 2010). The L4 nitrate data-set covered the whole 1998-2020 period, and despite several data-gaps, during most of this period it was sampled approximately with 5-7 day frequency. (ii) Another independent nitrate test data were obtained from The Scottish Coastal Observatory Dataset (<https://data.marine.gov.scot/dataset/scottish-coastal-observatory-dataset-1997-2020>, see also Bresnan et al. (2016); Hindson et al. (2018)), covering five locations near the coastline of Scotland (Loch Ewe, Scalloway, Scapa, St.Abbs, Stonehaven), providing time-series of differing lengths: from the longest, covering the 2008-2020 period (Scapa, Scalloway), to the shortest, covering the 2017-2020 period (St.Abbs). The measurements at those stations had typically 5-7 day frequency. The locations of the L4 and the five Scottish Stations are all marked in Fig.1. It should be noted that none of the L4 data and the data from the Scottish stations were included in the ICES data-set.

Finally, after validating the NN model, we have run it for the full 1998-2020 period across the whole Copernicus NWES reanalysis domain (see Fig.3), taking Copernicus reanalysis, river and ERA-5 atmospheric forcing inputs (both interpolated into the Copernicus reanalysis domain) and producing a gap-free bi-decadal, daily, 7 km resolution reconstruction of nitrate. This final data-set underpins the results from this study.

2.3 Skill metrics

We used the Relative Performance (RP) skill metric to compare the performance of the NN model from this study with the reanalysis:

$$RP(\text{NN}, \text{Rean}) = 100 \cdot \frac{(|\text{NN} - \text{Obs}| - |\text{Rean} - \text{Obs}|)}{\text{Obs}}. \quad (1)$$

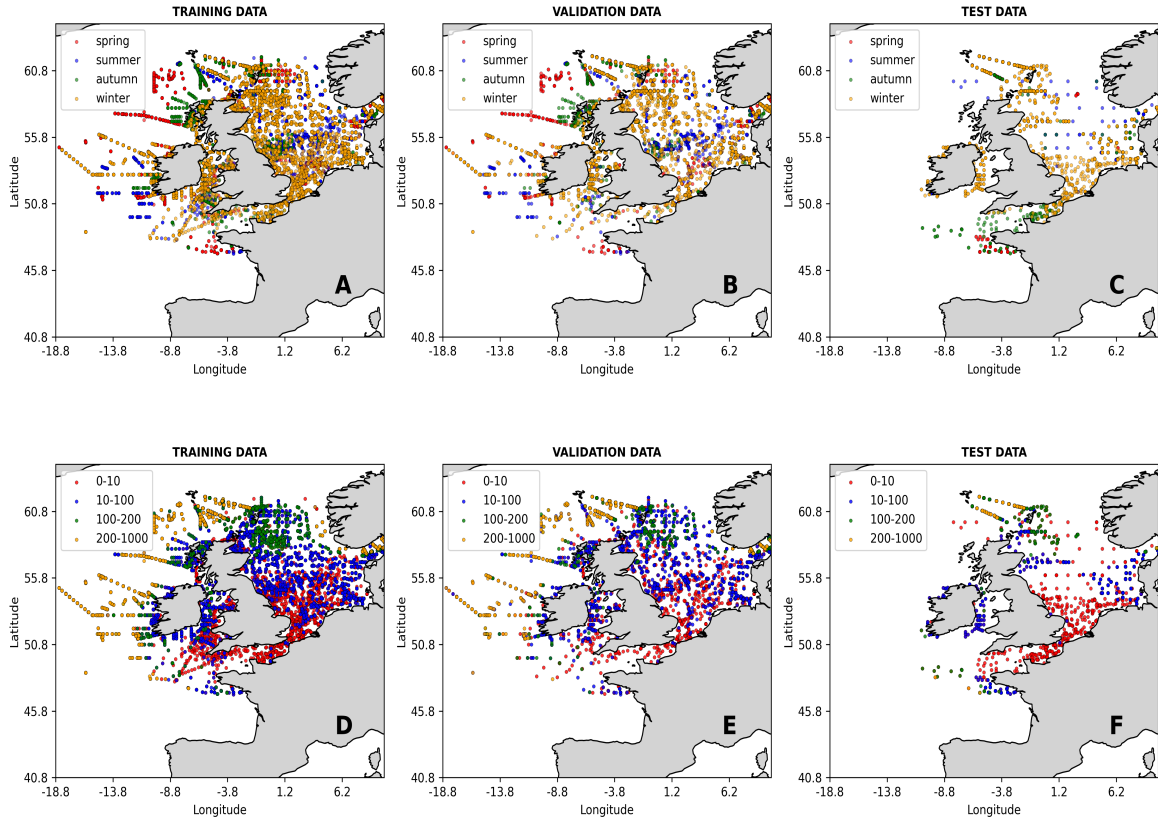


Figure 2. The locations of the ICES data used in this study split into training, validation and test data. The data-points are colored with respect to season when the measurement was taken (upper row) and by the depth range (in m) at which the measurement was taken (bottom row).

In Eq.1 “Obs” stands for observations, “NN” for predicted values by the NN model and “Rean” for reanalysis. Negative values of the RP metrics from Eq.1 indicate that NN model outperforms reanalysis and vice versa.

The bias between model and observations is defined as

$$170 \text{ Bias}(\text{Mod}, \text{Obs}) = \langle \text{Mod} - \text{Obs} \rangle, \quad (2)$$

where the averaging $\langle \dots \rangle$ is taken through all the available, matching, model and observational data.

The Bias-Corrected RMSE (BC-RMSE) is defined as RMSE after the bias as been subtracted from the model:

$$\text{BC-RMSE} = \sqrt{\langle (\text{Mod} - \text{Obs} - \text{Bias})^2 \rangle}. \quad (3)$$

Apart from the metrics from Eq.1-3, we have also used Pearson correlation. Since our tests have shown that the effective
 175 temporal resolution of the NN model (time-scale on which it performed best relative to the test data) is around 15 days (the

Table 1. The skill of the NN model in predicting nitrate compared with the Copernicus reanalysis (Kay et al., 2016). Skill is measured by bias (Eq.2, in $\mu\text{mol}/\text{m}^3$), Bias-Corrected Root Mean Squared Error (BC-RMSE, Eq.3, in $\mu\text{mol}/\text{m}^3$) and Pearson correlation (R). The rows represent different test data from ICES and the coastal stations. The last two rows show the skill of the NN model and the reanalysis to predict interannual, low-pass filtered time-series (for details see Fig.S3 of SI). For the five Scottish stations we show only the averaged result through all the stations.

Test data	NN predicted			Reanalysis		
	bias	BC-RMSE	R	bias	BC-RMSE	R
ICES	0.62	2.37	0.72	3.18	6.15	0.27
L4	-1.03	1.85	0.79	0.20	2.22	0.72
Scalloway	0.98	2.1	0.8	1.08	2.76	0.64
St.Abbs	-0.25	1.25	0.9	0.73	1.87	0.85
Scapa	1.52	1.52	0.86	0.33	2.13	0.69
Stonehaven	-0.53	0.97	0.95	0.09	2.04	0.78
Loch Ewe	1.57	0.92	0.93	0.57	1.1	0.89
L4 interannual	–	0.72	0.52	–	0.99	0.08
Scottish interannual	–	0.512	-0.144	–	0.756	0.204

tests are not shown here, but some insight is provided by Fig.S3 of SI), for each coastal station we have always compared the NN-predicted, reanalysis and observed data, only after they have been low-pass filtered on a 15-day scale.

3 Results and discussion

3.1 Model validation

180 Fig.4, Tab.1 and Fig.S4-S5 of SI demonstrate that the NN model shows a very good skill relative to the test data from ICES, L4 and the Scottish stations, and substantially outperforms the existing Copernicus reanalysis product for NWES nitrate. For example the bias measured by the ICES test data has been reduced by more than 80% relative to the reanalysis, the BC-RMSE has been reduced by more than 60% and the Pearson correlation has increased from 0.27 to 0.72. Comparison to the data from coastal stations shows less consistent improvement relative to the reanalysis, but the NN model still outperforms the reanalysis
185 at each of the locations.

Because the nitrate time series are dominated by the seasonal signal, it is important to explore whether the model skill extends beyond predicting the local nitrate seasonal (e.g. monthly) climatology. This is much harder to validate, as one needs long term time-series at specific locations, which are rare. We have looked at the data from the L4 station and five Scottish locations to analyse the ML model skill to capture interannual variability of nitrate. The results (shown in Tab.1 and Fig.S3 of
190 SI) are more mixed: at L4 station, which has from all the locations the longest time-record and richest data-set, the ML model

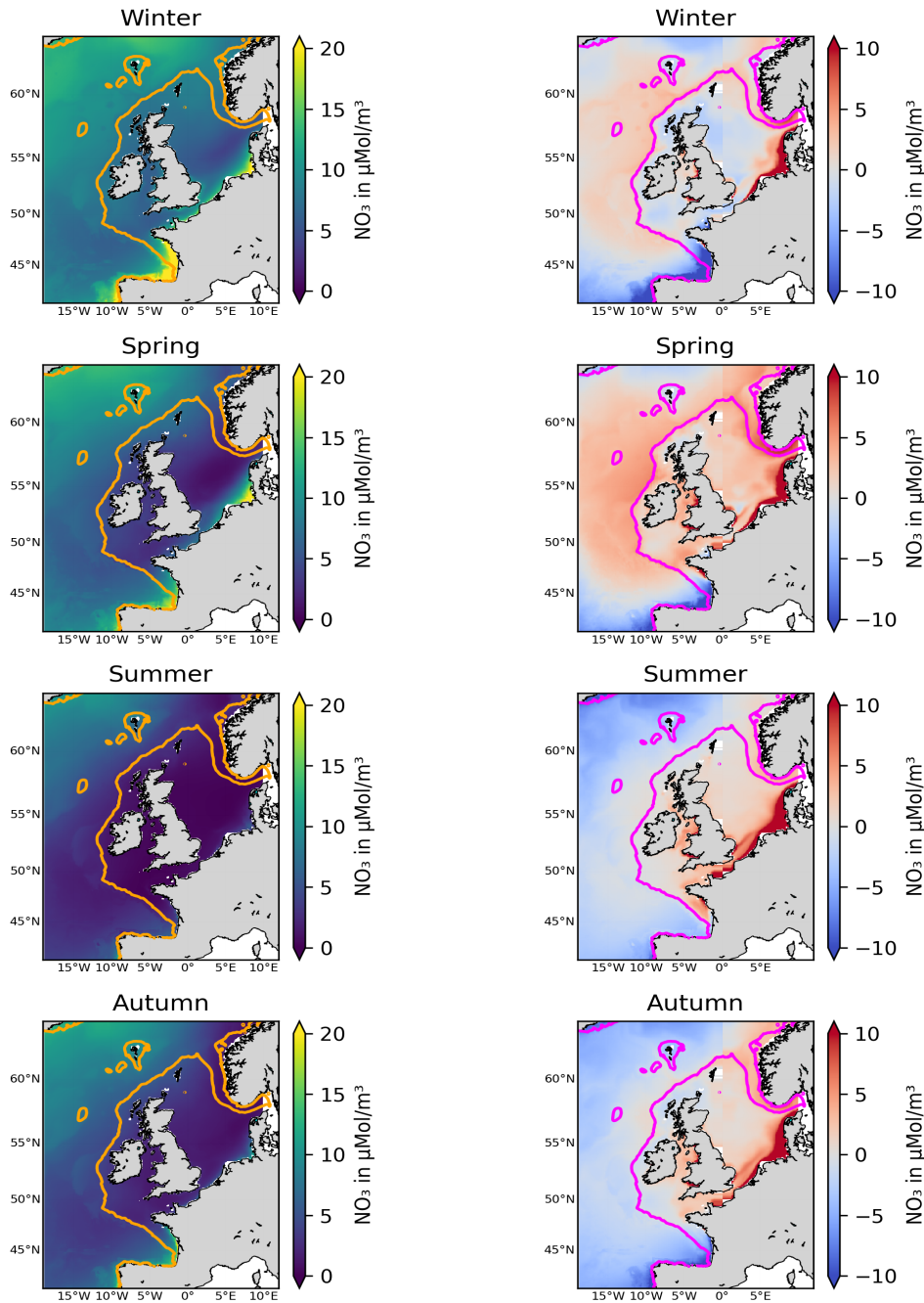


Figure 3. The left-hand panels show the NN-reconstructed 1998-2020 average surface nitrate concentrations for different annual seasons. The right-hand panels show the same averages for the relative bias of the Copernicus surface nitrate reanalysis (Kay et al., 2016) with respect to the NN-reconstructed data-set (reanalysis minus NN-reconstructed). The contours mark NWES (bathymetry < 200 m).

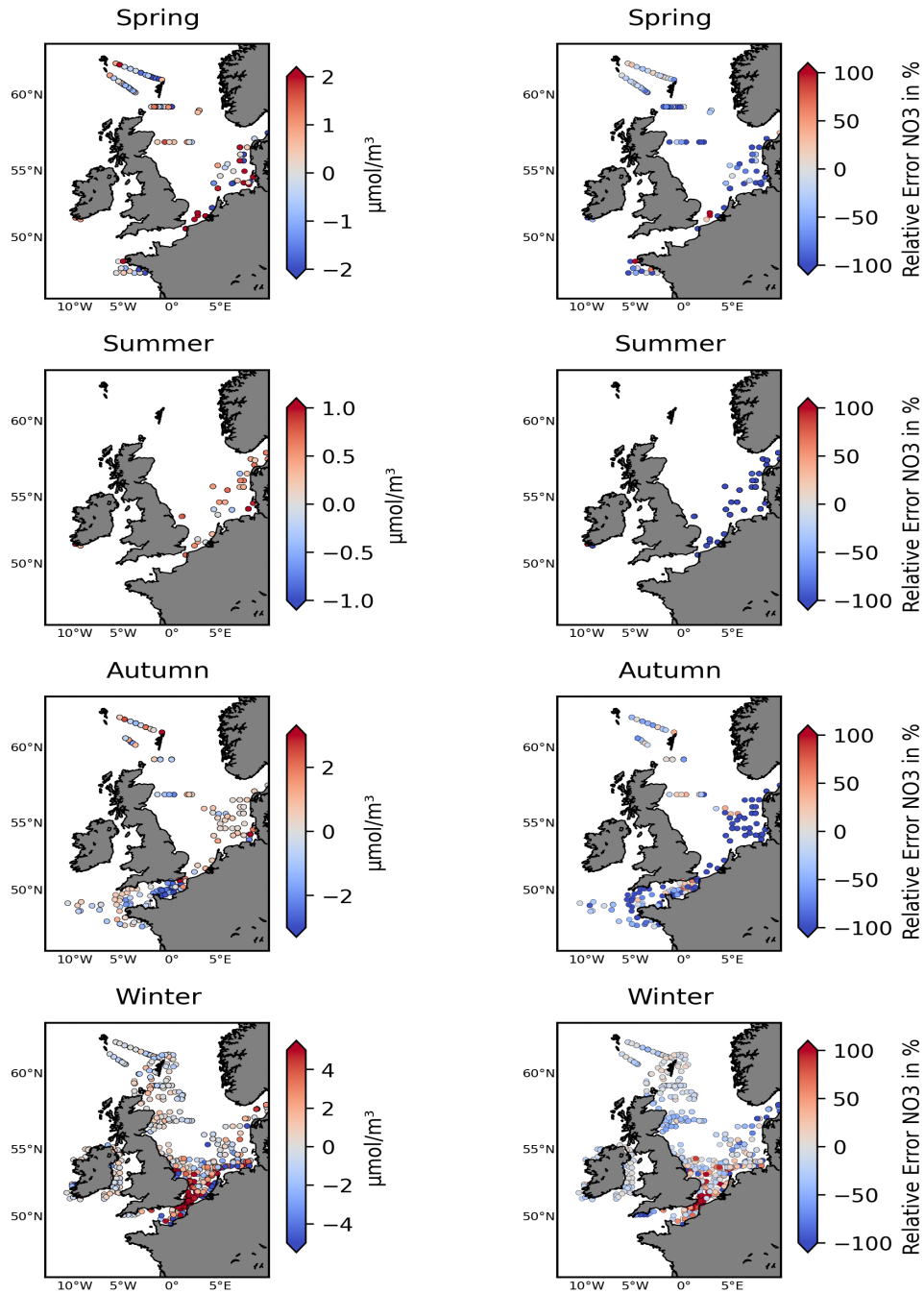


Figure 4. The left-hand panels show the ICES nitrate test data locations for different seasons, with the colorbar accounting for the NN-model skill (difference between predicted and observed nitrate: predicted - observed). The right-hand panels show the Relative Performance metrics between the NN model and the Copernicus reanalysis skill, as defined in Eq.1. It marks an NN-model improvement (blue), or degradation (red) relative to the reanalysis, when compared (in %) to the observed nitrate concentrations.

performs very well in predicting the inter-annual nitrate time-series. It is interesting that at the same location the reanalysis does a very poor job in doing the same (Tab.1). At the Scottish stations the ML model correctly captures the size of the interannual variability in nitrate, whereas struggles to capture the variability itself (the R metrics in Tab.1). It is however noteworthy that some of the time-series at the Scottish locations are relatively short (see Sec.2.1) and therefore not the most suitable for this type of analysis.

Finally, the test data selected from the ICES data-set are time-separated from the training and validation data, but were spatially located in largely overlapping regions (see Fig.2). It is therefore important to explore the possibility that, due to geographic proximity, some ML skill has been transferred from the training/validation data to the test data. This is done in Fig.S6 of SI, showing how the skill evolves as a function of spatial separation between the test data and the training/validation data. Although there is large variability in the skill, Fig.S6 shows no significant trend with spatial distance, indicating that the ML model skill does not decrease (even slightly improves) with the increase in spatial separation.

3.2 The bi-decadal nitrate product, the trends, variability and implications

Fig.3 shows the 1998-2020 seasonally averaged NWES nitrate concentrations. It is clear from the spatial nitrate distributions that the ML model does not capture sufficiently the ~ 7 km scale variability, including the exact NWES boundaries, but it does reasonably capture coarse resolution nitrate distributions (see Fig.5 for comparison with the World Ocean Atlas, WOA, product of Garcia et al. (2019)). Similarly, our analyses (including Fig.S3-S5 of SI) suggest that the effective temporal resolution of the NN product is ~ 15 -day, rather than daily. Fig.3 also provides seasonal comparison with the Copernicus reanalysis product, evaluating the significant reanalysis biases throughout the 1998-2020 period. The Copernicus reanalysis validation gives similar results to validation from Kay et al. (2016), who compared the reanalysis with the North Sea Biogeochemical 1960-2014 Climatology (Hinrichs et al., 2017).

The winter nitrate concentrations play an important role in pre-conditioning of the spring bloom, which largely drives the NWES biogeochemical seasonal cycles (Huisman et al., 1999; He et al., 2011). The winter total inorganic nitrogen is used by OSPAR, in combination with other parameters, using Common Procedure (OSPAR, 2005), as an important indicator for NWES eutrophication and next season's growth (Axe et al., 2017; Topcu and Brockmann, 2021). The hypothesis, that the intensity of spring phytoplankton bloom is directly related to the abundance of nutrients in the winter before the bloom, has been investigated here through nitrate. In Fig.6:A we have found only limited evidence for the relationship between the winter nitrate abundance and spring bloom intensity, i.e. statistically significant positive Pearson correlation has been found only in the western English Channel region, near the shelf-break in the Celtic Sea, around the Bay of Biscay and in the south-west of the model domain (accounting at most for 30-35% of explained variance). Fig.6:A also shows that these are regions where the inter-annual nitrate variability appears to be relatively large (10-20% of the winter average, Fig.6:B) and therefore capable to reveal stronger relationship with spring chlorophyll. For most of the domain, there is lack of clear correlation between inter-annual winter nitrate and spring chlorophyll, which could be explained by the fact that both are driven by the interannual variability in the atmosphere (Dutkiewicz et al., 2001; Follows and Dutkiewicz, 2001; Ueyama and Monger, 2005; Henson et al., 2006; Zhai et al., 2013). Increased winds can lead to more mixing and elevated surface nutrients, whilst dampening

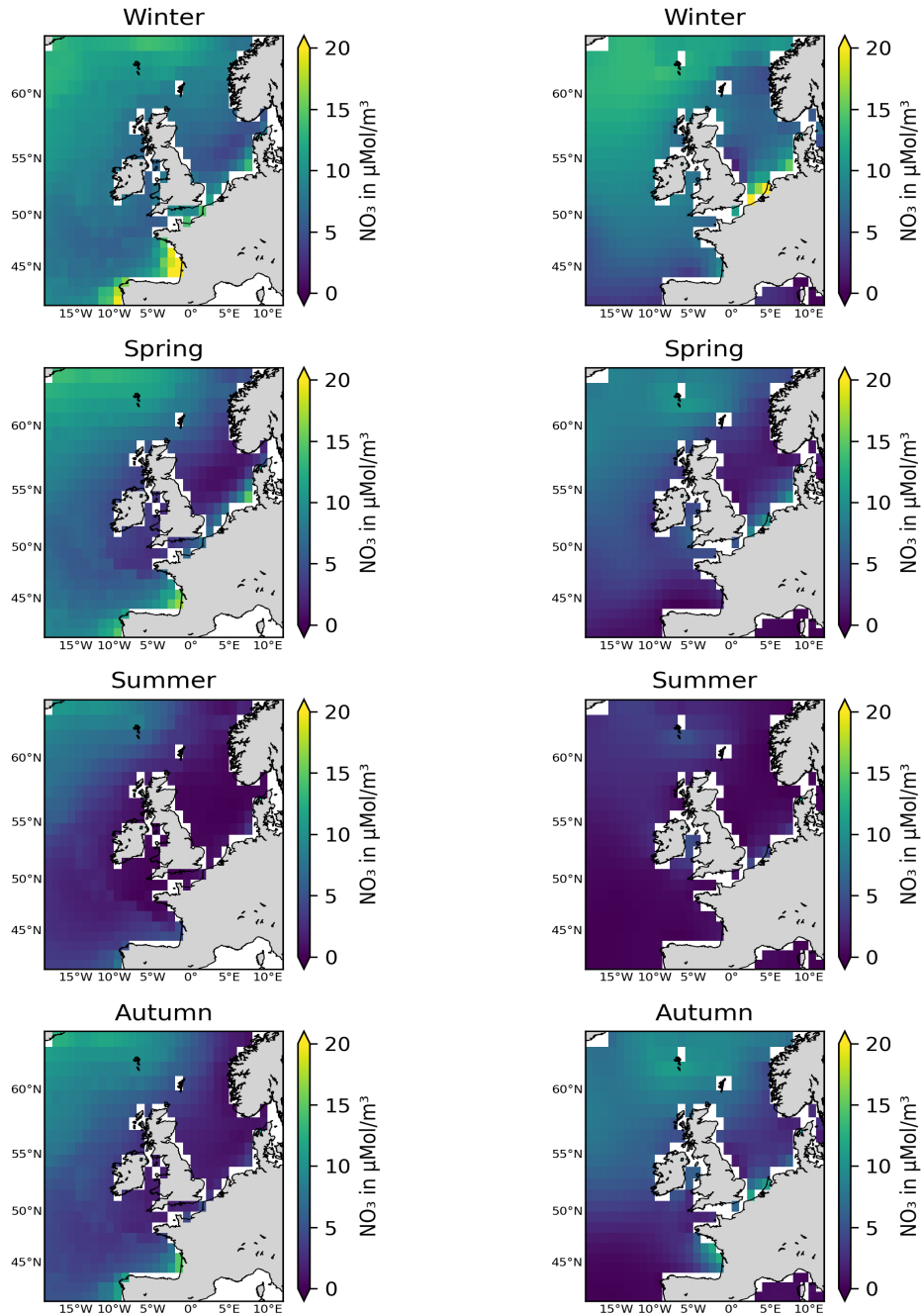


Figure 5. Seasonal comparison of the NN-predicted 1998-2020 surface nitrate averages with the World Ocean Atlas 2018 (WOA) product (Garcia et al., 2019). The NN-predicted data are coarse-grained on the WOA 0.25° spatial resolution scale. The focus is on spatial nitrate features, rather than nitrate concentration values, as the WOA atlas averages data for the whole 1990-2018 period, including highly eutrophic decades in the 1970's and 1980's.

225 blooms by transporting phytoplankton below the Sverdrup critical depth, as proposed by popular hypotheses explaining the North Atlantic spring blooms (Sverdrup, 1953; Huisman et al., 1999). Furthermore, there is lack of complete agreement on what are the dominant drivers of the spring bloom in the North Atlantic, and arguments have been raised supporting the view that blooms result much more from the internal ecosystem dynamics (e.g. zooplankton control over phytoplankton, Behrenfeld and Boss (2014)), compared to what was assumed by the traditional hypotheses focusing on physics.

230 In Fig.6:C we look at correlations between inter-annual time-series of summer nitrate and chlorophyll concentrations, indicating areas where phytoplankton is nitrate-limited (these are displayed by positive correlation). The Fig.6:C shows that chlorophyll is nitrate-limited mostly in the southern North Sea region, in the western English Channel, Bay of Biscay and the south-west of the domain. These are again the regions where the inter-annual fluctuations of summer nitrate are relatively large (Fig.6:D). The nitrate-limitation in these areas means they are vulnerable to eutrophication, if excess of nutrients is introduced
235 into the water. Indeed, it is re-assuring that the eutrophication-problem areas, as identified by the OSPAR NWES eutrophication status reports (such as south-eastern North Sea, coastal areas around Brittany, Axe et al. (2017)), fall under these vulnerable zones delimited in Fig.6:C. However, Fig.6:C includes also other regions, such as eastern coastline of Scotland, southern coast of Ireland and zones in the Irish Sea. Our results indicate that these additional regions could easily become problem-areas, if the policy and management of agriculture runoff became less effective.

240 Finally Fig.7, shows 1998-2020 trends in winter nitrate across the NWES domain. In most of the domain no statistically significant nitrate trends have been detected, but some small negative trends ($\sim 0.02 \mu\text{mol.m}^{-3}.\text{year}^{-1}$) were found in the Southern North Sea and the north-east region near the Norwegian trench. Somewhat larger ($\sim 0.08 \mu\text{mol.m}^{-3}.\text{year}^{-1}$) statistically significant negative trends have been found in specific locations of the Bay of Biscay. These results (e.g. from the Southern North Sea) are broadly consistent with what has been reported for this period in the recent OSPAR report (e.g. Axe
245 et al. (2017)). These small trends follow the smaller rates of reduction in the nitrate riverine inputs during the data period (1998-2020), compared to their large reduction in the 1980's and earlier 1990's (Duarte, 2009; Brockmann et al., 2018; Greenwood et al., 2019).

4 Conclusions

In this work we have demonstrated that, using sparse observations across the North-West European Shelf (NWES), machine
250 learning (ML) can provide a powerful tool to reconstruct spatially complete sea surface nitrate data-set over 22 year period. We have shown that the data-set has substantially better match-ups with independent test data than the existing NWES nitrate reanalysis. Using the newly developed product, we have identified nitrate-limited areas potentially vulnerable to eutrophication, addressed nitrate decadal trends, and tested how successfully winter nitrate can be used as a predictor of the phytoplankton spring bloom. There are many other potential scientific uses of the nitrate data-set, e.g. we propose to assimilate the nitrate
255 data into the NWES operational model, correcting the model significant nitrate biases, potentially improving its dynamics and its short-range forecasts. The model skill in simulating phytoplankton is known to quickly degrade with the forecast lead time (e.g. Kay et al. (2016); Skákala et al. (2018)) and biases in nitrate might be one of the leading factors in driving this.

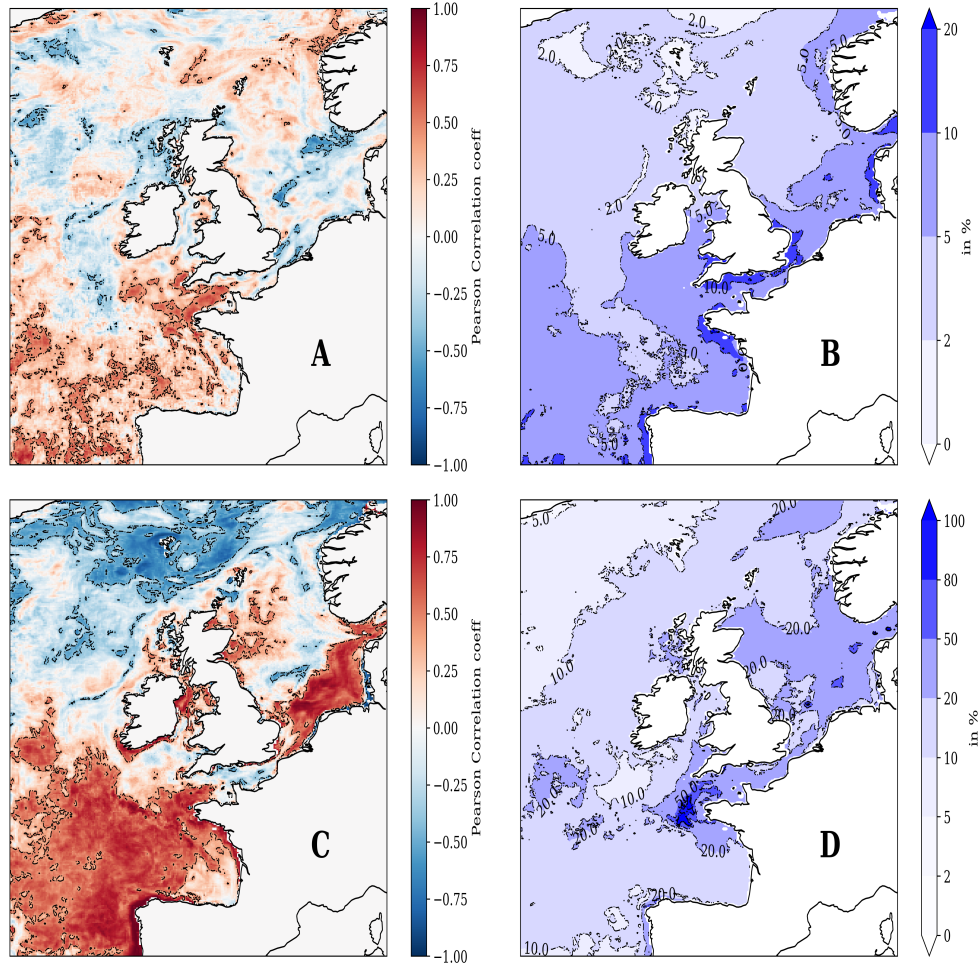


Figure 6. The upper left-hand panel (A) shows the Pearson correlation between the mean winter surface nitrate concentrations and the mean (following) spring surface chlorophyll concentration from the Copernicus reanalysis. The upper right-hand panel (B) shows the inter-annual variability for winter surface nitrate (across 1998-2020, measured by the standard deviation), relative to the 1998-2020 winter mean (in %). The bottom left panel (C) is similar to panel A, but showing the Pearson correlation between the summer surface nitrate and the summer surface total chlorophyll. The panel D is the same as the panel B, but showing inter-annual variability of surface nitrate in the summer, rather than winter. The dashed contours in panels A and C show regions where the correlation is statistically significant (p -value < 0.05).

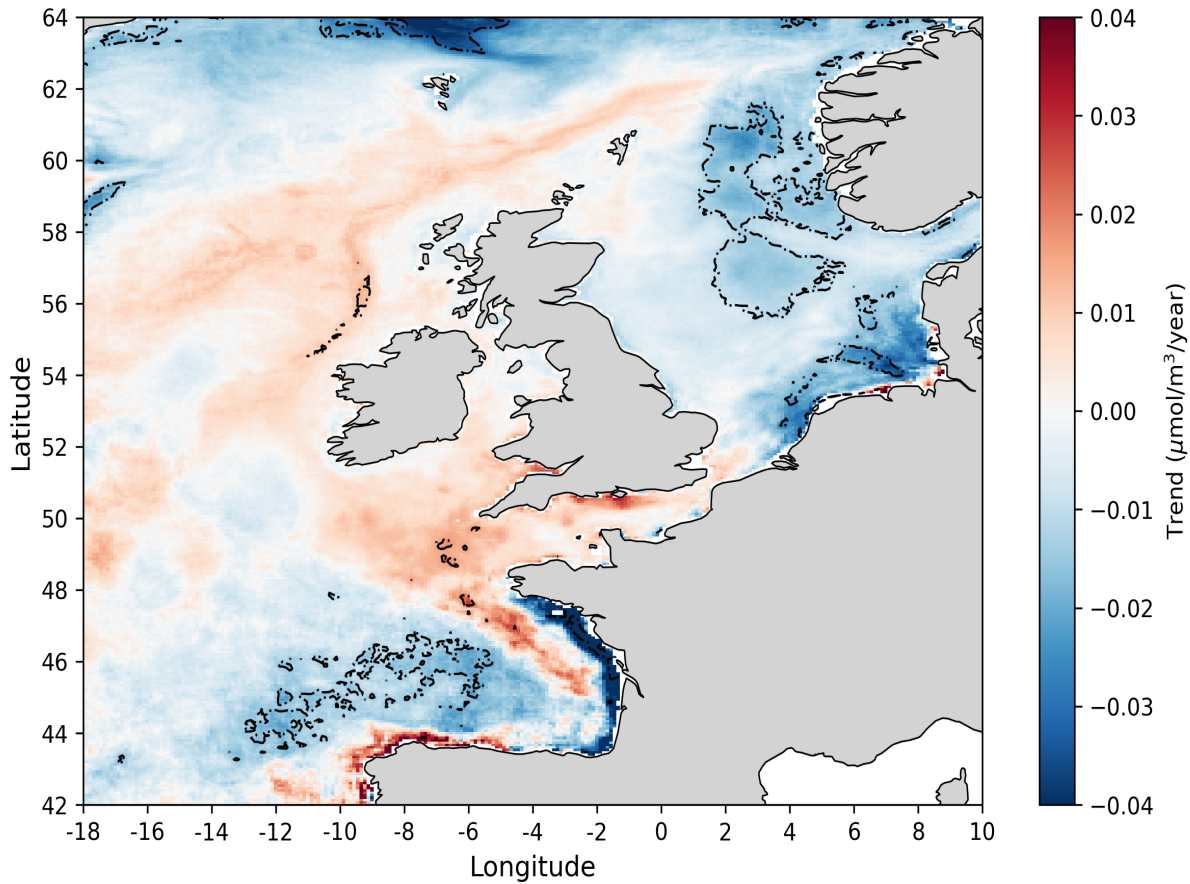


Figure 7. Linear trends at each spatial location in the annual nitrate 1998-2020 time-series. Dashed contours mark areas with statistically significant (p -value < 0.05) trends.

Several extensions of this work would be also desirable, such as utilizing ICES data for other biogeochemical indicators to produce ML-informed multi-variate data-sets across the whole NWES domain (these should include other nutrients and oxygen). ML could also identify valuable patterns of relationships across the multiple variables. Furthermore, the model developed here did not show very good skill in capturing high-frequency (daily) temporal variability, including extreme events. This might be due to processes providing ocean with memory significantly longer than the daily time-scale of the product. Representing ocean memory by the NN model might require using time-lagged input features, which could substantially inflate the size and the complexity of the model. Despite of that, including such features into the NN model should be considered in the future. Finally, ML tools designed to specifically capture extreme phenomena can be deployed in the future and extend the applicability of this work.

Code and data availability. The ML software is placed in <https://github.com/neccton-algo>. We used in this study the atmospheric ERA-5 product of the European Centre for Medium-range Weather Forecasting (ECMWF, <https://www.ecmwf.int/>), as well as river data that are stored on MonSOON HPC and can be obtained upon request. We have also used EU Copernicus reanalyses; the NWSHELF_MULTIYEAR_BGC_004_011 (<https://doi.org/10.48670/moi-00059>) product for biogeochemistry, and NWSHELF_MULTIYEAR-PHY-004_009 (<https://doi.org/10.48670/moi-00058>) for physics. We have used nitrate data-sets from the ICES portal (<https://doi.org/10.17895/ices.pub.8883>), from the Western Channel Observatory (https://www.westernchannelobservatory.org.uk/l4_nutrients.php), and the Scottish Coastal Observatory ([doi:10.7489/610-1](https://doi.org/10.7489/610-1), [10.7489/953-1](https://doi.org/10.7489/953-1), [10.7489/952-1](https://doi.org/10.7489/952-1), [10.7489/948-1](https://doi.org/10.7489/948-1), [10.7489/12138-1](https://doi.org/10.7489/12138-1))

Author contributions. DSB wrote all the code, processed the data, developed the ML model and run the experiments. JS provided conceptualisation, supervision and funding acquisition. Both authors wrote the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. This work was funded by the Horizon Europe project The New Copernicus Capability for Tropic Ocean Networks (NECCTON, grant agreement no.101081273). We also acknowledge support from the UK Natural Environment Research Council (NERC) single centre national capability programme – Climate Linked Atlantic Sector Science (CLASS, NE/R015953/1). The river data used here were prepared by Sonja van Leeuwen and Helen Powley as part of UK Shelf Sea Biogeochemistry programme (contract no.NE/K001876/1) of the NERC and the Department for Environment, Food and Rural Affairs (DEFRA). The riverine data contained also climatological values from the Global River Discharge Data Base and the Centre for Ecology and Hydrology (Young and Holt, 2007). We would like to thank Bee Bex for pointing us to the validation data from Scottish Coastal Stations. We would also like to thank Jerry Blackford, Gennadi Lessin, Yuri Artioli, Helen Powley, Julien Brajard and Dave Moffat for their valuable comments and discussions.

285 **References**

- Anderson, D. M., Cembella, A. D., and Hallegraeff, G. M.: Progress in understanding harmful algal blooms: paradigm shifts and new technologies for research, monitoring, and management, *Annual review of marine science*, 4, 143–176, 2012.
- Axe, P., Clausen, U., Leujak, W., Malcolm, S., and Harvey, E.: Eutrophication status of the OSPAR maritime area, *Third Integrated Report on the Eutrophication Status of the OSPAR Maritime Area*, 2017.
- 290 Behrenfeld, M. J. and Boss, E. S.: Resurrecting the ecological underpinnings of ocean plankton blooms, *Annual review of marine science*, 6, 167–194, 2014.
- Beman, J. M., Popp, B. N., and Francis, C. A.: Molecular and biogeochemical evidence for ammonia oxidation by marine Crenarchaeota in the Gulf of California, *The ISME Journal*, 2, 429–441, 2008.
- Board, O. S., Council, N. R., et al.: Clean coastal waters: understanding and reducing the effects of nutrient pollution, *National Academies Press*, 2000.
- 295 Borges, A., Schiettecatte, L.-S., Abril, G., Delille, B., and Gazeau, F.: Carbon dioxide in European coastal waters, *Estuarine, Coastal and Shelf Science*, 70, 375–387, 2006.
- Bresnan, E., Cook, K., Hindson, J., Hughes, S., Lacaze, J., Walsham, P., Webster, L., and Turrell, W.: The Scottish coastal observatory 1997–2013. Part 2-description of Scotland’s coastal waters, *Scottish Marine and Freshwater Science*, 7, 2016.
- 300 Brewin, R. J., Sathyendranath, S., Hirata, T., Lavender, S. J., Barciela, R. M., and Hardman-Mountford, N. J.: A three-component model of phytoplankton size class for the Atlantic Ocean, *Ecological Modelling*, 221, 1472–1483, 2010.
- Brewin, R. J., Ciavatta, S., Sathyendranath, S., Jackson, T., Tilstone, G., Curran, K., Airs, R. L., Cummings, D., Brotas, V., Organelli, E., et al.: Uncertainty in ocean-color estimates of chlorophyll for phytoplankton groups, *Frontiers in Marine Science*, 4, 104, 2017.
- Brockmann, U. and Eberlein, K.: River input of nutrients into the German Bight, in: *The role of freshwater outflow in coastal marine ecosystems*, pp. 231–240, Springer, 1986.
- 305 Brockmann, U., Topcu, D., Schütt, M., and Leujak, W.: Eutrophication assessment in the transit area German Bight (North Sea) 2006–2014–Stagnation and limitations, *Marine pollution bulletin*, 136, 68–78, 2018.
- Bruggeman, J. and Bolding, K.: A general framework for aquatic biogeochemical models, *Environmental modelling & software*, 61, 249–265, 2014.
- 310 Burson, A., Stomp, M., Akil, L., Brussaard, C. P., and Huisman, J.: Unbalanced reduction of nutrient loads has created an offshore gradient from phosphorus to nitrogen limitation in the North Sea, *Limnology and Oceanography*, 61, 869–888, <https://doi.org/10.1002/LNO.10257>, 2016.
- Butenschön, M., Clark, J., Aldridge, J. N., Allen, J. I., Artioli, Y., Blackford, J., Bruggeman, J., Cazenave, P., Ciavatta, S., Kay, S., et al.: ERSEM 15.06: a generic model for marine biogeochemistry and the ecosystem dynamics of the lower trophic levels, *Geoscientific Model Development*, 9, 1293–1339, 2016.
- 315 Diaz, R. J. and Rosenberg, R.: Spreading dead zones and consequences for marine ecosystems, *science*, 321, 926–929, 2008.
- Doney, S. C., Fabry, V. J., Feely, R. A., and Kleypas, J. A.: Ocean acidification: the other CO₂ problem, *Annual review of marine science*, 1, 169–192, 2009.
- Donlon, C. J., Martin, M., Stark, J., Roberts-Jones, J., Fiedler, E., and Wimmer, W.: The operational sea surface temperature and sea ice 320 analysis (OSTIA) system, *Remote Sensing of Environment*, 116, 140–158, 2012.

- Duarte, C. M.: Coastal eutrophication research: a new awareness, in: Eutrophication in Coastal Ecosystems: Towards better understanding and management strategies Selected Papers from the Second International Symposium on Research and Management of Eutrophication in Coastal Ecosystems, 20–23 June 2006, Nyborg, Denmark, pp. 263–269, Springer, 2009.
- 325 Duce, R. A., LaRoche, J., Altieri, K., Arrigo, K. R., Baker, A. R., Capone, D., Cornell, S., Dentener, F., Galloway, J., Ganeshram, R. S., et al.: Impacts of atmospheric anthropogenic nitrogen on the open ocean, *Science*, 320, 893–897, 2008.
- Dutkiewicz, S., Follows, M., Marshall, J., and Gregg, W. W.: Interannual variability of phytoplankton abundances in the North Atlantic, *Deep Sea Research Part II: topical studies in oceanography*, 48, 2323–2344, 2001.
- Follows, M. and Dutkiewicz, S.: Meteorological modulation of the North Atlantic spring bloom, *Deep Sea Research Part II: Topical Studies in Oceanography*, 49, 321–344, 2001.
- 330 Garcia, H., Weathers, K., Paver, C., Smolyar, I., Boyer, T., Locarnini, M., Zweng, M., Mishonov, A., Baranova, O., Seidov, D., et al.: World ocean atlas 2018. Vol. 4: Dissolved inorganic nutrients (phosphate, nitrate and nitrate+ nitrite, silicate), 2019.
- Good, S., Fiedler, E., Mao, C., Martin, M. J., Maycock, A., Reid, R., Roberts-Jones, J., Searle, T., Waters, J., While, J., et al.: The current configuration of the OSTIA system for operational production of foundation sea surface temperature and ice concentration analyses, *Remote Sensing*, 12, 720, 2020.
- 335 Greenwood, N., Devlin, M. J., Best, M., Fronkova, L., Graves, C. A., Milligan, A., Barry, J., and Van Leeuwen, S. M.: Utilizing eutrophication assessment directives from transitional to marine systems in the Thames Estuary and Liverpool Bay, UK, *Frontiers in Marine Science*, 6, 116, 2019.
- Harris, R.: The L4 time-series: the first 20 years, *Journal of Plankton Research*, 32, 577–583, 2010.
- He, R., Chen, K., Fennel, K., Gawarkiewicz, G., and McGillicuddy Jr, D.: Seasonal and interannual variability of physical and biological 340 dynamics at the shelfbreak front of the Middle Atlantic Bight: nutrient supply mechanisms, *Biogeosciences*, 8, 2935–2946, 2011.
- Henson, S. A., Robinson, I., Allen, J. T., and Waniek, J. J.: Effect of meteorological conditions on interannual variability in timing and magnitude of the spring bloom in the Irminger Basin, North Atlantic, *Deep Sea Research Part I: Oceanographic Research Papers*, 53, 1601–1615, 2006.
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., Simons, A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., Chiara, G. D., Dahlgren, P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C., Radnoti, G., de Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J. N.: The ERA5 global reanalysis, *Quarterly Journal of the Royal Meteorological Society*, 146, 1999–2049, <https://doi.org/10.1002/QJ.3803>, 2020.
- 350 Hill, R., Rinker, R., and Wilson, H. D.: Atmospheric nitrogen fixation by lightning, *Journal of Atmospheric Sciences*, 37, 179–192, 1980.
- Hindson, J., Berx, B., Hughes, S., Walsham, P., Machairpoulou, M., Bresnan, E., and Turrell, B.: The Scottish Coastal Observatory, *Bollettino di Geofisica*, p. 333, 2018.
- Hinrichs, I., Gouretski, V., Pätch, J., Emeis, K., and Stammer, D.: North sea biogeochemical climatology, 2017.
- Huisman, J., van Oostveen, P., and Weissing, F. J.: Critical depth and critical turbulence: two different mechanisms for the development of 355 phytoplankton blooms, *Limnology and oceanography*, 44, 1781–1787, 1999.
- Huthnance, J. M., Holt, J. T., and Wakelin, S. L.: Deep ocean exchange with west-European shelf seas, *Ocean Science*, 5, 621–634, 2009.
- Jahnke, R. A.: Global synthesis1, in: Carbon and nutrient fluxes in continental margins: A global synthesis, pp. 597–615, Springer, 2010.

- Jin, H., Song, Q., and Hu, X.: Auto-keras: An efficient neural architecture search system, in: Proceedings of the 25th ACM SIGKDD international conference on knowledge discovery & data mining, pp. 1946–1956, 2019.
- 360 Kay, S., McEwan, R., and Ford, D.: North West European Shelf Production Centre NWSHELF_MULTIYEAR_BIO_004_011, CMEMS Report, 3, 21, 2016.
- Lenhart, H.-J., Mills, D. K., Baretta-Bekker, H., van Leeuwen, S. M., van der Molen, J., Baretta, J. W., Blaas, M., Desmit, X., Kühn, W., Lacroix, G., Los, H. J., Ménesguen, A., Neves, R., Proctor, R., Ruardij, P., Skogen, M. D., Vanhoutte-Brunier, A., Villars, M. T., and Wakelin, S. L.: Predicting the consequences of nutrient reduction on the eutrophication status of the North Sea, *Journal of Marine Systems*, 81, 148–170, <https://doi.org/https://doi.org/10.1016/j.jmarsys.2009.12.014>, contributions from Advances in Marine Ecosystem Modelling Research II 23-26 June 2008, Plymouth, UK, 2010.
- 365 Madec, G., Bourdallé-Badie, R., Bouttier, P.-A., Bricaud, C., Bruciaferri, D., Calvert, D., Chanut, J., Clementi, E., Coward, A., Delrosso, D., et al.: NEMO ocean engine, 2017.
- Nazari-Sharabian, M., Ahmad, S., and Karakouzian, M.: Climate change and eutrophication: a short review, *Engineering, Technology and Applied Science Research*, 8, 3668, 2018.
- 370 Noxon, J.: Atmospheric nitrogen fixation by lightning, *Geophysical Research Letters*, 3, 463–465, 1976.
- OSPAR, C.: Common procedure for the identification of the eutrophication status of the OSPAR maritime area, OSPAR Commission, 3, 2005.
- Pauly, D., Christensen, V., Guénette, S., Pitcher, T. J., Sumaila, U. R., Walters, C. J., Watson, R., and Zeller, D.: Towards sustainability in world fisheries, *Nature*, 418, 689–695, 2002.
- 375 Postgate, J. R.: Nitrogen fixation, Cambridge University Press, 1998.
- Rabalais, N. N., Turner, R. E., and Wiseman Jr, W. J.: Gulf of Mexico hypoxia, aka “The dead zone”, *Annual Review of ecology and Systematics*, 33, 235–263, 2002.
- Rabalais, N. N., Turner, R. E., Díaz, R. J., and Justić, D.: Global change and eutrophication of coastal waters, *ICES Journal of Marine Science*, 66, 1528–1537, 2009.
- 380 Radach, G.: Ecosystem functioning in the German Bight under continental nutrient inputs by rivers, *Estuaries*, 15, 477–496, 1992.
- Ryther, J. H. and Dunstan, W. M.: Nitrogen, phosphorus, and eutrophication in the coastal marine environment, *Science*, 171, 1008–1013, 1971.
- Skákala, J., Ford, D., Brewin, R. J., McEwan, R., Kay, S., Taylor, B., de Mora, L., and Ciavatta, S.: The assimilation of phytoplankton functional types for operational forecasting in the northwest European shelf, *Journal of Geophysical Research: Oceans*, 123, 5230–5247, 2018.
- 385 Skákala, J., Ford, D., Bruggeman, J., Hull, T., Kaiser, J., King, R. R., Loveday, B., Palmer, M. R., Smyth, T., Williams, C. A., et al.: Towards a multi-platform assimilative system for North Sea biogeochemistry, *Journal of Geophysical Research: Oceans*, 126, e2020JC016 649, 2021.
- 390 Skákala, J., Bruggeman, J., Ford, D., Wakelin, S., Akpınar, A., Hull, T., Kaiser, J., Loveday, B. R., O’Dea, E., Williams, C. A., et al.: The impact of ocean biogeochemistry on physics and its consequences for modelling shelf seas, *Ocean Modelling*, 172, 101 976, 2022.
- Sverdrup, H. U.: On conditions for the vernal blooming of phytoplankton, *J. Cons. Int. Explor. Mer*, 18, 287–295, 1953.
- Tett, P., Droop, M. R., and Heaney, S. I.: The Redfield Ratio and Phytoplankton Growth Rate, *Journal of the Marine Biological Association of the United Kingdom*, 65, 487–504, <https://doi.org/10.1017/S0025315400050566>, 1985.

- 395 Topcu, D. and Brockmann, U.: Consistency of thresholds for eutrophication assessments, examples and recommendations, *Environmental Monitoring and Assessment*, 193, 1–15, 2021.
- Ueyama, R. and Monger, B. C.: Wind-induced modulation of seasonal phytoplankton blooms in the North Atlantic derived from satellite observations, *Limnology and Oceanography*, 50, 1820–1829, 2005.
- Voss, M., Bange, H. W., Dippner, J. W., Middelburg, J. J., Montoya, J. P., and Ward, B.: The marine nitrogen cycle: recent discoveries, 400 uncertainties and the potential relevance of climate change, *Philosophical Transactions of the Royal Society B: Biological Sciences*, 368, 20130 121, 2013.
- Withers, P. J., Neal, C., Jarvie, H. P., and Doody, D. G.: Agriculture and eutrophication: where do we go from here?, *Sustainability*, 6, 5853–5875, 2014.
- Young, E. and Holt, J.: Prediction and analysis of long-term variability of temperature and salinity in the Irish Sea, *Journal of Geophysical Research: Oceans*, 112, 2007. 405
- Zhai, L., Platt, T., Tang, C., Sathyendranath, S., and Walne, A.: The response of phytoplankton to climate variability associated with the North Atlantic Oscillation, *Deep Sea Research Part II: Topical Studies in Oceanography*, 93, 159–168, 2013.