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DATA DESCRIPTOR

Carbon dioxide fluxes in Alpine grasslands at the Nivolet Plain, Gran Paradiso National Park, Italy 2017–2023

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We introduce a georeferenced dataset of Net Ecosystem Exchange (NEE), Ecosystem Respiration (ER) and meteo-climatic variables (air and soil temperature, air relative humidity, soil volumetric water content, pressure, and solar irradiance) collected at the Nivolet Plain in Gran Paradiso National Park (GNPN), western Italian Alps, from 2017 to 2023. NEE and ER are derived by measuring the temporal variation of CO₂ concentration obtained by the enclosed chamber method. We used a customised portable non-steady-state dynamic flux chamber, paired with an InfraRed Gas Analyser (IRGA) and a portable weather station, measuring CO₂ fluxes at a number of points (around 20 per site and per day) within five different sites during the snow-free season (June to October). Sites are located within the same hydrological basin and have different geological substrates: carbonate rocks (site CARB), gneiss (GNE), glacial deposits (GLA, EC), alluvial sediments (AL). This dataset provides relevant and often missing information on high-altitude mountain ecosystems and enables new comparisons with other similar sites, modelling developments and validation of remote sensing data.

Background & Summary

Earth's changing climate is significantly affecting mountain ecosystems¹. Temperature rise and modification of precipitation patterns lead to glacier retreat, reduction of snow cover, alteration of the water cycle, and impacts on living organisms and biogeochemical cycles. In particular, climate change can affect the structure and functioning of mountain ecosystems, particularly for what concerns the natural carbon cycle. Previous research indicated that natural grasslands act as a net carbon sink^{2–4}. However, the carbon fluxes and carbon storage capacity of these ecosystems are likely to change in response to climate warming, particularly in high-mountain areas that are more susceptible to temperature rise^{5,6}. Quantifying the carbon fluxes at the soil-vegetation-atmosphere interface in high-mountain ecosystems, and simultaneously measuring meteo-climatic and environmental variables, is an essential source of information for investigating what are the main drivers of carbon fluxes and understanding the response of CO₂ fluxes to climate change.

To this aim, in 2017 the Institute of Geosciences and Earth Resources of the National Research Council of Italy (IGG-CNR) established an Alpine Critical Zone Observatory (CZO@NIVOLET) in the north-western Italian Alps (Nivolet Plain, Gran Paradiso National Park).

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The dataset presented here is the result of data collection using portable non-steady-state flux chambers and weather stations. The dataset contains measurements collected at individual points within the study sites approximately every 10–15 days during the snow-free period in seven years of fieldwork, from 2017 to 2023.

The non-steady-state flux chamber is a classical method used for estimating gas fluxes, in particular greenhouse gases such as CO₂ and CH₄, from different types of interfaces, including bare soil^{7,8} and natural ecosystems^{9,10}.

Part of the data presented here were already published as average values of point-measurements for each site and each sampling date, and are freely available in the IGG-CNR-CZO community of the Zenodo repository^{11–13}.

In a related research article titled “Drivers of carbon fluxes in Alpine tundra: a comparison of three empirical model approaches”¹⁴ multi-regression models were developed for Gross Primary Production (GPP, defined as GPP = NEE - ER) and Ecosystem Respiration (ER) using average values for each site and each sampling date from years 2017, 2018, and 2019. Further investigations, based on the above-mentioned average values and additional data from CZO@NIVOLET, have also been discussed in “Carbon dioxide exchanges in an alpine tundra ecosystem (Gran Paradiso National Park, Italy): A comparison of results from different measurement and modelling approaches”¹⁵ and “Spatial and temporal variability of carbon dioxide fluxes in the Alpine Critical Zone: The case of the Nivolet Plain, Gran Paradiso National Park, Italy”¹⁶.

In this manuscript, we present and make freely available the complete dataset of point-measurements, which were previously analysed only as averages over site and sampling date. Moreover, this dataset includes data from the 2023 field campaigns, which have never been used nor published in any form before.

This dataset enables new modelling and analysis efforts by the scientific community. It can be used for spatio-temporal analysis of CO₂ fluxes in Alpine ecosystems, for comparisons with CO₂ fluxes from other environments, and for validating models developed by using remote sensing data. It can also be used for diagnostic purposes in the analysis of the dependence of CO₂ fluxes on climate drivers.

In addition to sharing the complete dataset with the research community, this manuscript provides a comprehensive description of the CZO@NIVOLET site’s methodology for data collection and processing using the portable flux chamber method. This description encompasses each step of the process, from the instruments’ calibration at the dedicated laboratory in a controlled environment to the calculation of CO₂ fluxes, reported as μmolCO₂ m⁻² s⁻¹. Furthermore, we emphasise that the CZO@NIVOLET site remains actively investigated. This manuscript then provides guidance to the understanding and utilisation of present and forthcoming data generated in this study site which will be as well updated within the IGG-CNR-CZO community of the Zenodo repository.

Methods

Study site. The CZO@NIVOLET was installed in 2017 and is located within the boundaries of the Gran Paradiso National Park (GPNP). It is part of the Critical Zone Exploration Network (CZEN, <https://www.czen.org/content/nivolet-czo>), a global network investigating processes in the Critical Zone, which is defined as the dynamic living skin of the Earth that extends from the top of the vegetative canopy through the soil and down to fresh bedrock and the bottom of the groundwater¹⁷. This research site also belongs to both the European eLTER (<https://elter-ri.eu/elter-ri>) and ICOS ERIC (<https://www.icos-cp.eu>) Research Infrastructures (RI).

The GPNP was established in 1922 for the preservation of the Alpine ibex (*Capra ibex*) and the conservation of high-altitude mountain ecosystems. Encompassing an area of 720 km², the park features a wide range of ecosystems, including lower elevation Alpine woods, as well as high-altitude grasslands and Alpine tundra, rock cliffs, and glaciers above the treeline. The Nivolet Plain (Fig. 1) is a glacial valley that ranges in elevation from approximately 2300 m a.s.l. in the northeast to around 2700 m a.s.l. in the southwest.

The underlying bedrock is composed of gneisses, dolostones and marbles from the Gran Paradiso Massif, as well as calcschists with serpentinites and metabasites from the Piedmont-Ligurian zone¹⁸.

Daily records of precipitation (mm), minimum, maximum, and mean temperature (°C) from the Lago Agnel weather station are freely available (CC BY-NC-SA 4.0) at Arpa Piemonte portal (https://www.arpa.piemonte.it/rischi_naturali/snippets_arpa_graphs/dati_giornalieri_meteo/?statid=PIE-001073-900-1996-10-10¶m=P). According to such data, over the time span 2017–2023, the average daily minimum temperature from June to October was 6.11 °C, the average daily maximum temperature was 13.0 °C, and the average daily precipitation was 2.9 mm. During winter the soil is typically covered with a thick layer of snow.

The Nivolet Plain is home to Alpine natural grasslands that support a diverse array of species within the *Caricion curvulae* climax vegetation community¹⁹. Dominant species found in the grasslands include *Carex curvula* All., *Alopecurus gerardi* Vill., *Gnaphalium supinum* L., and *Leontodon helveticus* Mérat. In the investigation sites, also *Geum montanum*, *Trifolium alpinum*, *Pulsatilla alpina*, and *Silene acaulis* are commonly found. The plants in these high-altitude grasslands experience rapid development from late June to late October, with canopy heights reaching a maximum of 0.2 metres.

During summer, grasslands are grazed by both domestic and wild ungulates. Wild ungulates (ibex and chamois) are censused every year. In 2022, the population density in the GPNP was counted to be 2687 ibex individuals and 6346 chamois individuals (GPNP, unpublished data, see also²⁰), while there are no quantitative census data on roe deer, red deer, and wild boar. These latter, however, are typically found at lower altitudes than those considered in our study. Regarding domestic ungulates, from the beginning of July until mid-September, approximately 110 cows along with around 20 sheep, and goats are brought for grazing in this area¹⁹. Grazing is conducted in a controlled manner, with animals predominantly grazing in areas adjacent to the barn (located at 45°29'13.7″N 7°08'27.6″E) and throughout the lower regions of the Nivolet valley (see <https://www.pastoralpe.it/homepage/> for more information).



Fig. 1 Location of the CZO@NIVOLET. The Nivolet Plain ($45^{\circ}28'42.96''\text{N}$ $7^{\circ}08'31.92''\text{E}$) is located in the north-western Italian Alps. The image was acquired by Landsat/Copernicus, sourced and modified from Google Earth.

The five measurement sites at the Nivolet Plain are located within the same hydrological basin, and each site has an area between ~ 500 to ~ 900 square metres. Three of the five selected sites are on the orographic left flank bordering the Nivolet Plain: one on carbonate rocks (site named CARB in the dataset, at 2750–2760 m a.s.l.) and two on glacial deposits (site named GLA, at 2740–2750 m a.s.l. and site named EC, at 2750–2760 m a.s.l.). One site lies on the orographic right flank of the Plain, on soils developed on gneiss (site named GNE, at about 2580–2600 m a.s.l.). One site is on alluvial soil at the Plain floor (site named AL, 2740–2750 m a.s.l.). The location of the five sites is shown in Fig. 2. Mean coordinates of the five study sites are reported in Table 1.

In 2020, D'Amico *et al.*²¹ published the soil types map of the Aosta Valley, encompassing the region that includes our study area. Besides, our research group conducted soil profile samplings in locations proximate to and with similar geological and geomorphological attributes of the five study sites. Some physical and chemical characteristics of soil profiles are briefly described here^{22,23}.

Since 2020, data on aboveground vegetation biomass at the EC site have become available. These values, calculated as averages from individual samples, are detailed in Table 2.

Flux chamber measurements. In the summer of 2017, surveys were carried out to select measurement sites, assess instrumental setups, and determine the CO_2 flux ranges essential for laboratory calibration of portable flux chamber systems. This initial phase was characterised by few measurement campaigns. Starting from 2018, the frequency of these campaigns increased, establishing a regular schedule of measurements approximately every 10–15 days throughout the vegetative season. The final instrumentation setup is shown in Fig. 3.

NEE and ER were measured using the non-steady state dynamic flux chamber method. The chamber, placed over the vegetated soil, isolates a volume of air where the concentration of CO_2 increases or decreases according to the dominant process at the soil-vegetation-atmosphere interface. In the presence of sunlight, if photosynthesis captures CO_2 faster than its release due to respiration, the CO_2 concentration inside the chamber decreases. If respiration is dominant over photosynthesis, the CO_2 concentration increases. CO_2 concentration inside the flux chamber is measured over a specific time interval; the flux is computed by interpolating the curve of CO_2 concentration versus time, as explained further in the text and discussed in detail here²⁴.

At each point, a stainless-steel collar was inserted for about 1 cm into the soil a few minutes before the measurement, assuring no leakage. Before placing the flux chamber, RGB (Red, Green, Blue) images were taken from a nadir perspective, aiming at monitoring the vegetation within the collar area. These images are freely available in the IGG-CNR-CZO community of the Zenodo repository²⁵.

The flux chamber was then placed on the collar to isolate a confined air volume (headspace). This created a closed system where the CO_2 concentration inside the chamber changes during the measurement because

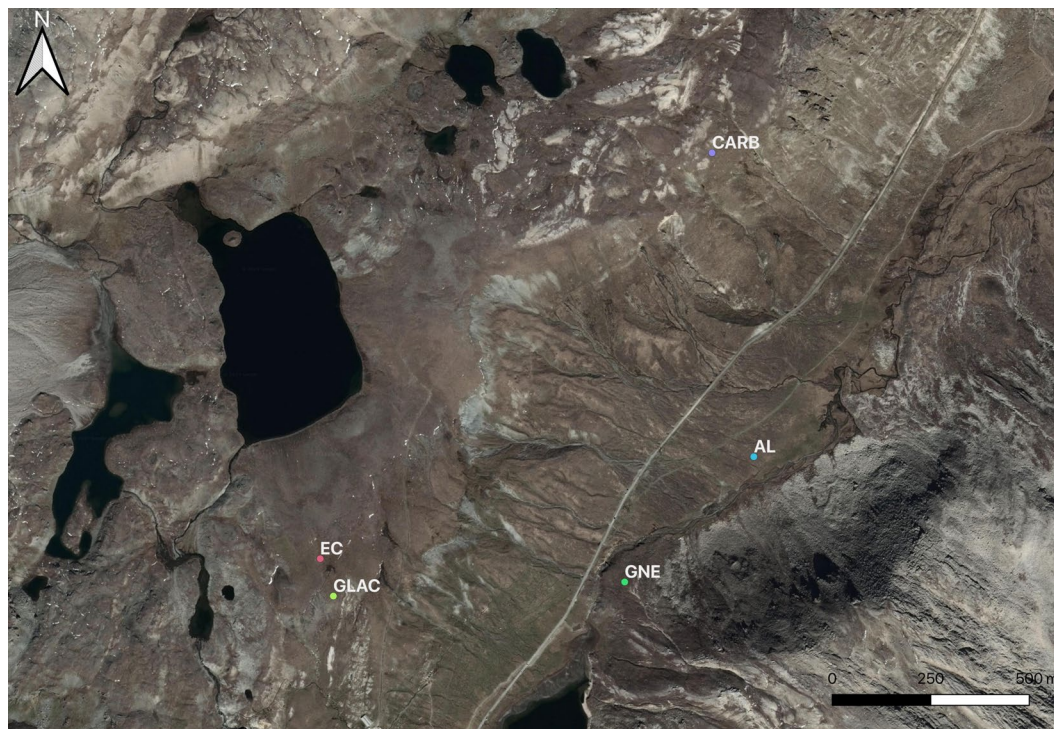


Fig. 2 Location of the five measurement sites. Soils developed on carbonate rocks (CARB), gneiss (GNE), glacial deposits (GLAC & EC) and alluvial sediments (AL). Made with Qgis software (v.3.16 Hannover, QGIS Development Team, 2021, www.qgis.org) (Map data ©2015 Google).

| Site | MEAN Lon | MEAN Lat | MEAN Elevation |
|------|----------|----------|--------------------|
| AL | 7.15345 | 45.49324 | 2490–2500 m a.s.l. |
| CARB | 7.15210 | 45.50016 | 2750–2760 m a.s.l. |
| EC | 7.13937 | 45.49091 | 2750–2760 m a.s.l. |
| GLAC | 7.13979 | 45.49006 | 2740–2750 m a.s.l. |
| GNE | 7.14926 | 45.49038 | 2580–2600 m a.s.l. |

Table 1. Mean coordinates of the five study sites. Mean coordinates were calculated by averaging the latitude and longitude of individual measurement points. First column: site name; second and third columns: mean longitude and mean latitude (WGS84); fourth column: mean elevation.

| Year | Biomass | N. of Samples | Date of Sampling |
|------|---------|---------------|------------------|
| 2020 | 0.23 | 20 | 08/08/2020 |
| 2021 | 0.197 | 16 | 14/07/2021 |
| 2023 | 0.14 | 15 | 26/07/2023 |

Table 2. Mean biomass values for aboveground vegetation at EC site. Mean biomass values were calculated by averaging individual vegetation samples. The biomass samples were obtained by harvesting 0.25 m x 0.25 m aboveground vegetation plots at the maximum growing season. The samples were then dried to constant weight at 60 °C for 48 hours and weighed. These measurements were conducted at the ICOS associated station Nivolet (ICOS code IT-Niv), which corresponds to the EC site. The procedure was conducted in accordance with the guidelines outlined here³² for ICOS associated stations and followed a modified protocol derived from the ICOS procedures described here³³. These, along with other data, are freely accessible here³⁴ upon registration on the ICOS data portal (ICOS CCBY4 Data Licence). First column: sampling year in yyyy format; second column: mean biomass, expressed in kilograms of dry matter per square metre (kgDM m^{-2}); third column: total number of biomass samples collected for averaging; fourth column: date of sampling in dd/mm/yyyy format.

of CO_2 absorption by plants through photosynthesis and/or emission through respiration by autotrophs (i.e., plants) and heterotrophs (i.e., microbial communities in the soil).

Air from the headspace of the chamber was pumped at a constant flow rate of 3 l/min into an Infrared Gas Analyzer (IRGA, either model LI-840 or LI-850 $\text{CO}_2/\text{H}_2\text{O}$ Analyzer; LI-COR Biosciences, Lincoln, NE, USA)

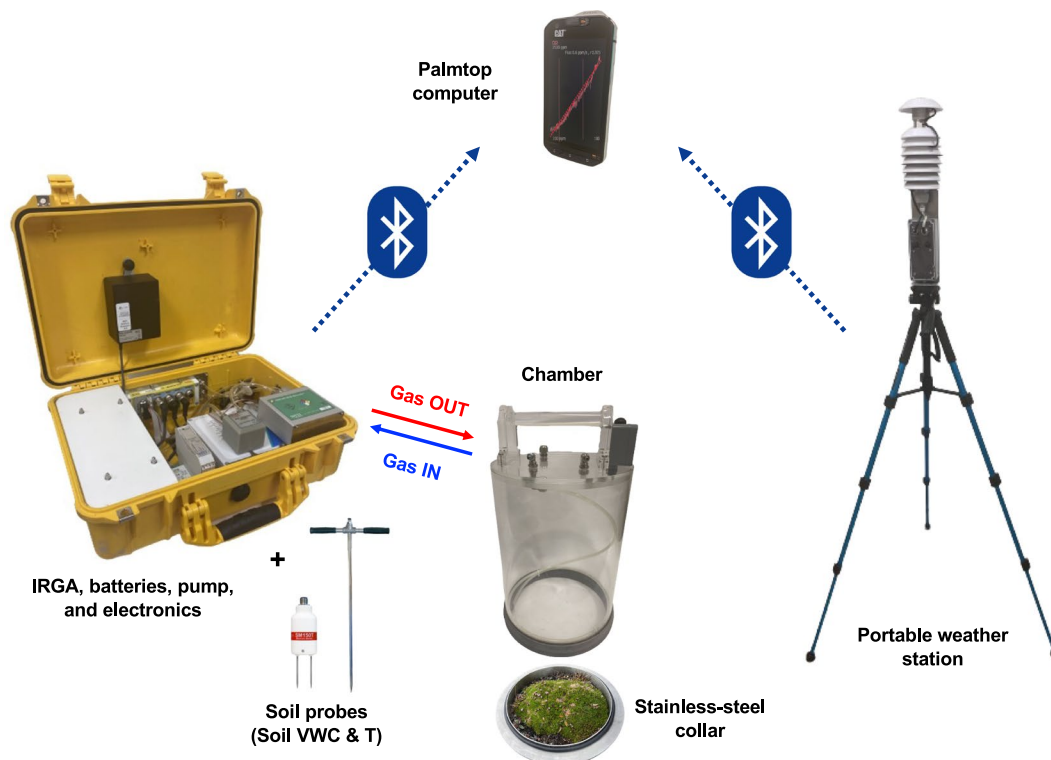


Fig. 3 Portable instrumentation setup. The yellow case on the left contains the IRGA (Infrared Gas Analyzer), batteries, pump, and electronics. The IRGA is connected to the flux chamber through two RILSAN[®] tubing pipes (gas IN and gas OUT), each measuring 1.8 m in length and having an internal diameter of 4 mm and an external diameter of 6 mm. The gas sampling line is protected by two types of filters: (1) a 50 mm diameter PTFE membrane filter with a pore size of 0.45 μm , and (2) a 25 mm diameter PTFE membrane filter with a pore size of 0.2 μm . These filters are permeable to gases and water vapour but are impermeable to liquid water and dust particles. The soil volumetric water content is recorded using a TDR (Time-Domain Reflectometry) soil sensor. The soil temperature is recorded using a Pt100 soil thermometer. All data are recorded at 1 Hz during the measurement. Air relative humidity, air temperature, and solar irradiance are measured by a portable weather station (thermohygrometer and pyranometer) mounted on a tripod at a height of 1.5 metres above the ground (on the right). An Android device (palm-top computer) connected via Bluetooth serves as an interface for managing the measurement, displaying, and storing the data.

through a 1.8-metres-long RILSAN[®] tubing. The sampled air was then reinjected into the chamber. The reinjection tube ended with a 0.40-metres-long coiled and pierced RILSAN[®] tube, which ensured good mixing of the reinjected air sample within the chamber. Before and after each measurement, the entire apparatus (including the chamber, tubing, and IRGA) was vented until the ambient CO_2 concentration was recorded and its concentration was stable for a few seconds.

The CO_2 concentration inside the flux chamber was measured for about 90 seconds. CO_2 concentration versus time was recorded at 1 Hz frequency using the custom Android app FluxManager2 (West Systems S.r.l., freely available on Google Play Store) which was installed on a palm-top computer connected to the instrument via Bluetooth. Upon completion of the measurements, a text file containing all the data, including meteorological and environmental variables (see below), was generated in the internal memory of the Android device.

The concentration curve was interpolated linearly over a period of about 60 seconds to calculate the rate of change of CO_2 concentration over time (ppm s^{-1}). The interpolation was done using the custom FluxRevision software (West Systems S.r.l.). The initial 10–15 seconds (cleaning time), and the final, potentially non-linear part of the curve were excluded from the interpolation.

Measurements were conducted at various times throughout the day, ranging from 10:00 to 18:00⁹, and covered different meteorological conditions, in order to capture the natural meteorological variability²⁶. For each measurement campaign and for each site, measurements were replicated at 15–20 different points within the site, randomly chosen to sample the small-scale flux variability. Previous analysis has shown that a minimum of 15 measurement points is generally sufficient to represent the spatial variability at these sites¹⁴.

Figure 4 shows a typical measurement cycle, which included two consecutive measurements at each point: the first was performed under ambient light, using the transparent chamber to estimate NEE, while the second measurement was performed using the same chamber shaded with a cloth to estimate ER in the absence of photosynthesis. A similar procedure was applied in previous works on similar environments^{26,27}. This process was repeated at all points, requiring approximately 2 hours to cover an entire site. Notice that in 2020, owing to the restrictions imposed by the pandemics, the number of measurement campaigns had to be much reduced.



Fig. 4 Measurement procedure. The standard measurement cycle consists of two consecutive measurements at each point. The first measurement is conducted under natural light conditions using the transparent chamber (left) to determine the NEE. The second measurement is performed using the shaded chamber (right) to determine the ER. All individuals in figures have provided explicit consent for their images to be openly published.

NEE and ER fluxes in $\mu\text{molCO}_2 \text{ m}^{-2} \text{ s}^{-1}$ were estimated from the slope of the linear regression of headspace CO_2 concentration over time (ppm s^{-1}) using a laboratory calibration curve that relates pre-determined CO_2 fluxes (in the range of fluxes expected in the field) with the corresponding measured slopes of the CO_2 vs time linear regression (see the section “Technical Validation”).

Mean values and variability of Net Ecosystem Exchange (NEE) and Ecosystem Respiration (ER) measured at site GNE (2017–2023) are illustrated in Fig. 5, as an example of the data from one of the five sites. Part of the NEE data discussed here have been compared with the flux estimates provided by an eddy covariance tower located at the EC site, belonging to the FLUXNET network as ICOS-Associated ecosystemic station since 2022 (IT-Niv, ref to: https://meta.icos-cp.eu/resources/stations/ES_IT-Niv). The results of the comparison indicated that the site-average of the individual NEE point measurements at the EC site were consistent with the NEE estimates provided by the eddy covariance method for the same time and date¹⁵.

Meteo-climatic variables. The optimised version of the portable weather station, shown in Fig. 3, was employed starting from 2018.

During the CO_2 flux measurements, FluxManager2 simultaneously recorded air temperature, atmospheric pressure, air relative humidity, solar irradiance, soil temperature, and soil volumetric water content (1 Hz acquisition).

Air relative humidity, air temperature, and solar irradiance were measured using LSI LASTEM thermohygrometers model DMA672.1 sheltered from direct solar radiation and LSI LASTEM pyranometers model DPA053A mounted on a portable tripod at a height of 1.5 metres above the ground²⁸. The atmospheric pressure was recorded using digital barometers placed inside the flux chambers.

Soil temperature and soil volumetric water content were measured using Pt100 thermometers and Delta-T SM150T soil moisture sensors at depths of approximately 10 cm for the soil temperature and in the range of 0–5 cm for the soil moisture. The measurements were taken at about 20 cm from the collar on undisturbed soil, specifically without removing the organic layer (layer O). To account for small scale variability of soil moisture, soil volumetric water content values were also taken inside the collar area before the measurements of CO_2 concentration to assess the moisture range (at least 3 measurements), then the probe was placed outside the collar, at a point where the soil water content was in the range of values measured inside the collar.

To ensure accuracy, these sensors were tested at CNR laboratories before and after each measurement season and calibrated in accredited laboratories every two years. The specifications for the sensors and probes are provided in Table 3.

FluxManager2. The FluxManager2 Android application (West Systems S.r.l.) is installed on a palmtop computer provided with Bluetooth; it is used to manage the instrumentation, sensors and probes and for displaying and recording the data.

FluxManager2 Android app is freely available on Google Play Store.

FluxRevision. The FluxRevision software (West Systems S.r.l.) allows users to interpolate the CO_2 concentration curve and calculate their slope and R^2 , using files created with FluxManager2. The software leaves the possibility to choose the linear interpolation interval.

FluxRevision is freely available for download from the West Systems website (<https://www.westsystems.com/instruments/download/>).

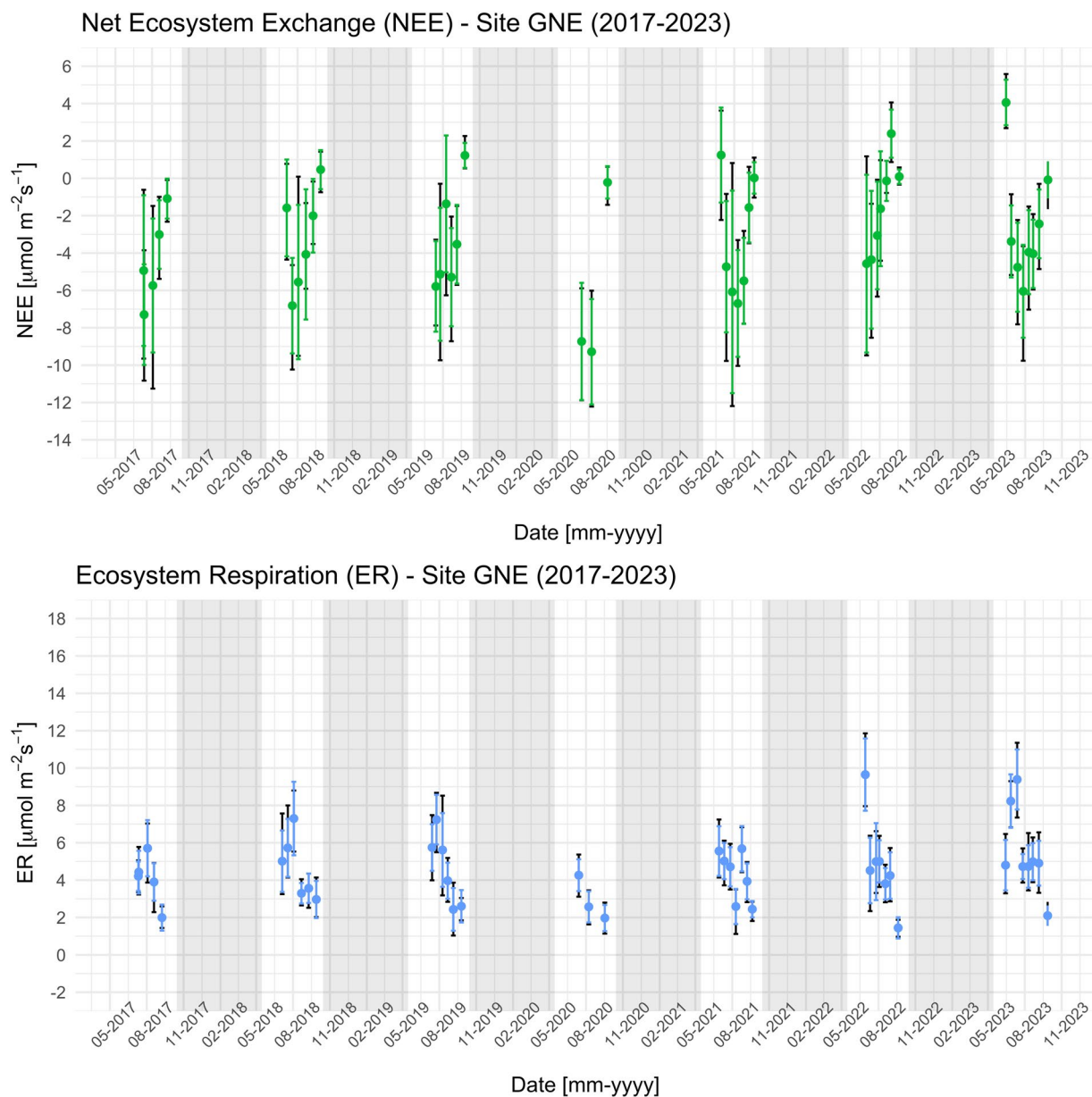


Fig. 5 Net Ecosystem Exchange (NEE, top) and Ecosystem Respiration (ER, bottom) measured at site GNE (2017–2023). The coloured dots represent the mean values, while the dark arrows indicate the 10th and 90th quantiles. Coloured bars depict the intervals of 1 standard deviation (σ).

Data Records

The dataset provided with this manuscript is organised as a comma-separated text file (.csv) and is available at the IGG-CNR-CZO Community page in the Zenodo repository²⁹.

Fields are separated by semicolons and NA indicates values that are Not Available or were discarded after data quality control (ref. to the following section “Technical Validation”). Each record includes all the values of the variables recorded at each single measurement point.

Sign convention is the following: the flux from the atmosphere *to* the soil/ecosystem (e.g., photosynthetic CO_2 uptake, GPP) is *negative*, whereas the flux *from* the soil/ecosystem (ER) to the atmosphere is *positive*. Thus, $\text{NEE} = \text{GPP} + \text{ER}$ can be either positive or negative. NEE and ER fluxes are reported in $\mu\text{molCO}_2 \text{ m}^{-2} \text{ s}^{-1}$.

Names/acronyms used in the dataset and their description are listed in Table 4. Meteo-climatic variables recorded during the measurement of NEE or during the measurement of ER bring the suffix NEE or ER respectively (i.e., Pressure_NEE = atmospheric pressure recorded during the measurement of Net Ecosystem Exchange).

A comprehensive workflow showing all the steps performed from data acquisition to the final dataset is reported in Fig. 6.

| Sensor/probe & variable measured | Model & Manufacturer | Range | Accuracy | Expanded measurement uncertainty |
|--|---|-----------------------------|---|---|
| Thermometer (AirT) | DMA672.1 LSI LASTEM Srl | [−50°:100°] | 0.1 °C (@0°C) | 0.2 °C |
| Hygrometer (Air_RH) | DMA672.1 LSI LASTEM Srl | [0:100%] | ±1% (@5:95%) | 2.0% |
| Pyranometer (SolarRad) | DPA053A LSI LASTEM Srl | [0:2000 W m ^{−2}] | ±1% (100:1000 W m ^{−2}) | 2.40% [W m ^{−2}] |
| Platinum Resistance Thermometer Pt100 RTD (SoilT) | Pt100 Industrial Sensor Probe, Class B RS PRO | [−50°:100°] | 0.05 °C | 0.055 °C |
| Time-domain reflectometry (TDR) soil probe (SoilVWC) | SM150T sensor Delta-T Devices Ltd. | [0:100%] | ±3.0% vol over 0%–70% vol, and at 0–60 °C | ±3.0% vol over 0%–70% vol, and at 0–60 °C |
| Digital barometer (Pressure) | MPL3115A2 NXP Semiconductors / Freescale | [500:1100 hPa] | ±0.5 hPa | ±2.5 hPa |

Table 3. Characteristics of the sensors and probes used to measure meteo-climatic variables.

| Variable name | Unit | Description |
|----------------|---|--|
| DATE | dd/mm/yy | Campaign date in dd/mm/yy format |
| YEAR | yyyy | Campaign year in yyyy format |
| TIME.UTC+2_NEE | H:M:S | Time of NEE measurement in UTC + 02 |
| TIME.UTC+2_ER | H:M:S | Time of ER measurement in UTC + 02 |
| Site | | Measurement site (AL, CARB, GLAC, GNE, EC) |
| LONG_(E) | | Longitude East, as recorded by the integrated palmtop computer GPS (WGS84 spatial reference system) |
| LAT_(N) | | Latitude North, as recorded by the integrated palmtop computer GPS (WGS84 spatial reference system) |
| ELEVATION | m | Altitude above sea level (m) |
| SolarRad_NEE | W m ^{−2} | Solar irradiance in [W m ^{−2}] during NEE measurement |
| AirT_NEE | °C | Air temperature in [°C] during NEE measurement |
| Air_RH_NEE | % | Air relative humidity in [%] during NEE measurement |
| Pressure_NEE | hPa | Atmospheric pressure in [hPa] during NEE measurement |
| SoilT_NEE | °C | Soil temperature in [°C] at 10 cm depth during NEE measurement |
| SoilVWC_NEE | % | Soil volumetric water content in the range of 0–5 cm depth in [%] during NEE measurement |
| FLUX_NEE | μmolCO ₂ m ^{−2} s ^{−1} | NEE CO ₂ flux obtained through a laboratory calibration curve that allowed to convert the temporal variation of the CO ₂ concentration inside the chamber [ppm s ^{−1}] into the CO ₂ flux [μmolCO ₂ m ^{−2} s ^{−1}]. |
| SLOPE_NEE | ppm s ^{−1} | Slope of the linear regression of the CO ₂ concentration vs time curve [ppm s ^{−1}] |
| R2_SLOPE_NEE | | R-squared of the linear regression of the CO ₂ concentration vs time curve [ppm s ^{−1}] |
| SolarRad_ER | W m ^{−2} | Solar irradiance in [W m ^{−2}] during ER measurement |
| AirT_ER | °C | Air temperature in [°C] during ER measurement |
| Air_RH_ER | % | Air relative humidity in [%] during ER measurement |
| Pressure_ER | hPa | Atmospheric pressure in [hPa] during ER measurement |
| SoilT_ER | °C | Soil temperature in [°C] at 10 cm depth during ER measurement |
| SoilVWC_ER | % | Soil volumetric water content in the range of 0–5 cm depth in [%] during ER measurement |
| FLUX_ER | μmolCO ₂ m ^{−2} s ^{−1} | ER CO ₂ flux obtained through a laboratory calibration curve that allowed to convert the temporal variation of the CO ₂ concentration inside the chamber [ppm s ^{−1}] into the CO ₂ flux [μmolCO ₂ m ^{−2} s ^{−1}]. |
| SLOPE_ER | ppm s ^{−1} | Slope of the linear regression of the CO ₂ concentration vs time curve [ppm s ^{−1}] |
| R2_SLOPE_ER | | R-squared of the linear regression of the CO ₂ concentration vs time curve [ppm s ^{−1}] |

Table 4. Description of the dataset variables.

Technical Validation

Before and after each measurement season, we tested and calibrated the instrumental equipment (including flux chamber, pump, IRGA, connecting tubes, and portable weather stations) to ensure proper functioning and performance. We use a calibration curve that is specific to the instrumental setup and is determined on a case-by-case basis to account for any variations or changes in the equipment over time.

Flux chamber calibration (Fig. 7) is conducted under controlled environmental conditions in the laboratory using reference CO₂ mass flow rates obtained from a high-precision thermal Mass Flow Controller (MFC) specifically designed for gases (red-y smart controller GSC, Vögtlin Instruments GmbH), and high-precision CO₂

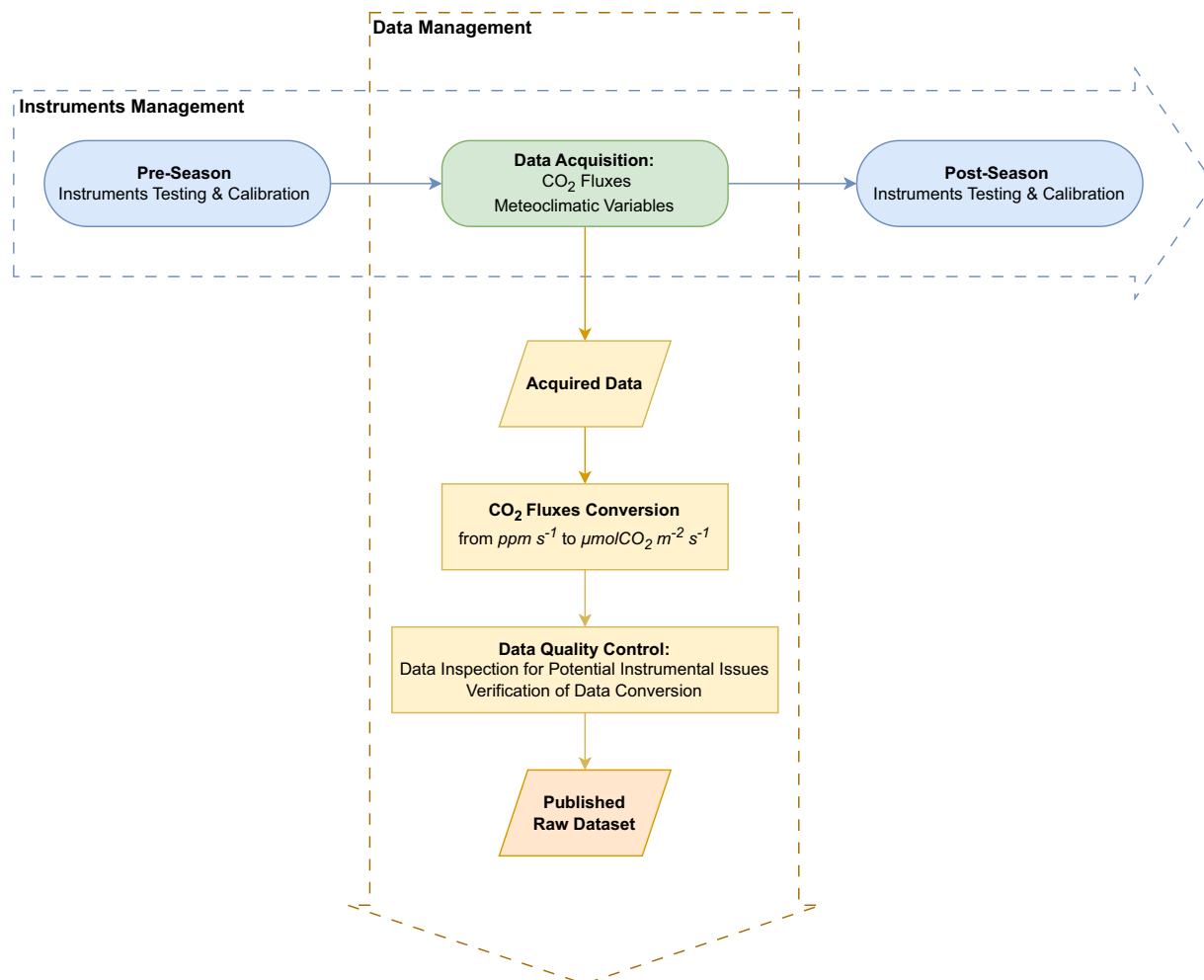


Fig. 6 Data Management (DM) and Instruments Management (IM) workflow. The DM workflow illustrates each step from data acquisition to the final product.

mixtures with certified concentrations. Two different high-precision CO₂ mixtures were used: 1) CO₂ 2.00%mol, CH₄ 1.00%mol, and N₂; 2) CO₂ 1.00%mol, CH₄ 500 ppm mol, and N₂. The calibration of the measurement apparatus is essential for reducing the uncertainty of CO₂ flux estimates. Our research group has been involved in investigating the uncertainty associated with CO₂ flux measurements, with a focus on very low fluxes, resulting in the publication of the article titled “*Non-steady-state closed dynamic chamber to measure soil CO₂ respiration: A protocol to reduce uncertainty*”²². The calibration process follows the same measurement procedures used in the field and the reference CO₂ mass flow rates were chosen to cover the range of fluxes expected in the field^{30,31} (but not exceeding two orders of magnitude²²). To test the reproducibility of the measurements, we perform 5 to 8 replicates at each predetermined CO₂ flux.

The laboratory tests indicated that the devices achieved good reproducibility for data acquisition times of 90 seconds. However, for very low fluxes (close to detection limit), it was necessary to increase the acquisition time up to 120–150 seconds to obtain reliable results.

Figure 8 is an example of a calibration curve. It shows the intercept of the linear interpolation of CO₂ concentration vs time obtained with the flux chamber (in ppm s⁻¹) versus the predetermined CO₂ fluxes (in cc min⁻¹).

The calibration curve is used in the conversion of the CO₂ fluxes measured in the field. Initially, CO₂ fluxes are corrected for the ratio between atmospheric pressure and air temperature recorded during the measurement, and those recorded in the laboratory when the calibration curve was obtained. Then, from the equation of the calibration curve, the CO₂ fluxes are initially converted in cc min⁻¹ - which is the measurement unit of the predetermined CO₂ used in the calibration curve - and then in μmolCO₂ s⁻¹. Finally, the obtained values are divided by the collar area (0.036 m²) to obtain the CO₂ fluxes in μmolCO₂ m⁻² s⁻¹.

In addition to calibration, the IRGA were checked periodically to ensure proper operation by performing the following tests:

1. Verifying the zero CO₂. It is verified by adding a CO₂ scrubber to the air inlet of the IRGA and by using a zero CO₂ cylinder in the laboratory (i.e., pure N₂). The CO₂ scrubber is used to reduce any atmospheric CO₂ contamination to zero and ensure accurate readings.

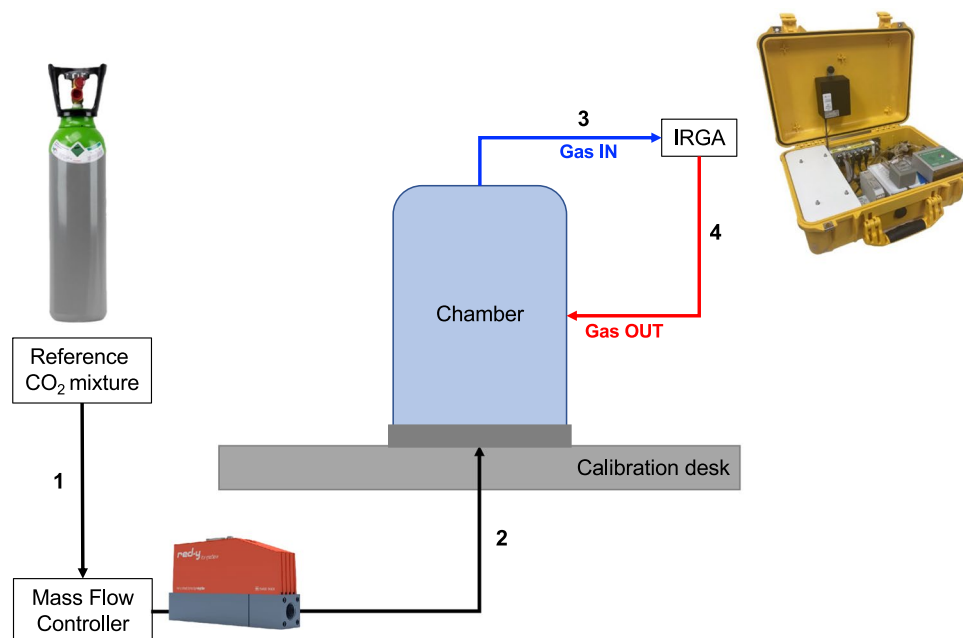


Fig. 7 Scheme of the calibration setup. A high-precision thermal mass flow controller is used to set a constant CO₂ mass flow (1), which is then routed inside the flux chamber through a hole in a rubber-covered desk (calibration desk) that simulates the soil surface (2). The CO₂ mass flow then enters the flux chamber, and the air from the headspace of the chamber is pumped at a constant flow rate of 3 l/min into the IRGA (3). The IRGA is used to measure the concentration of CO₂ in the air sample. Finally, the air sample is re-injected into the flux chamber (4).

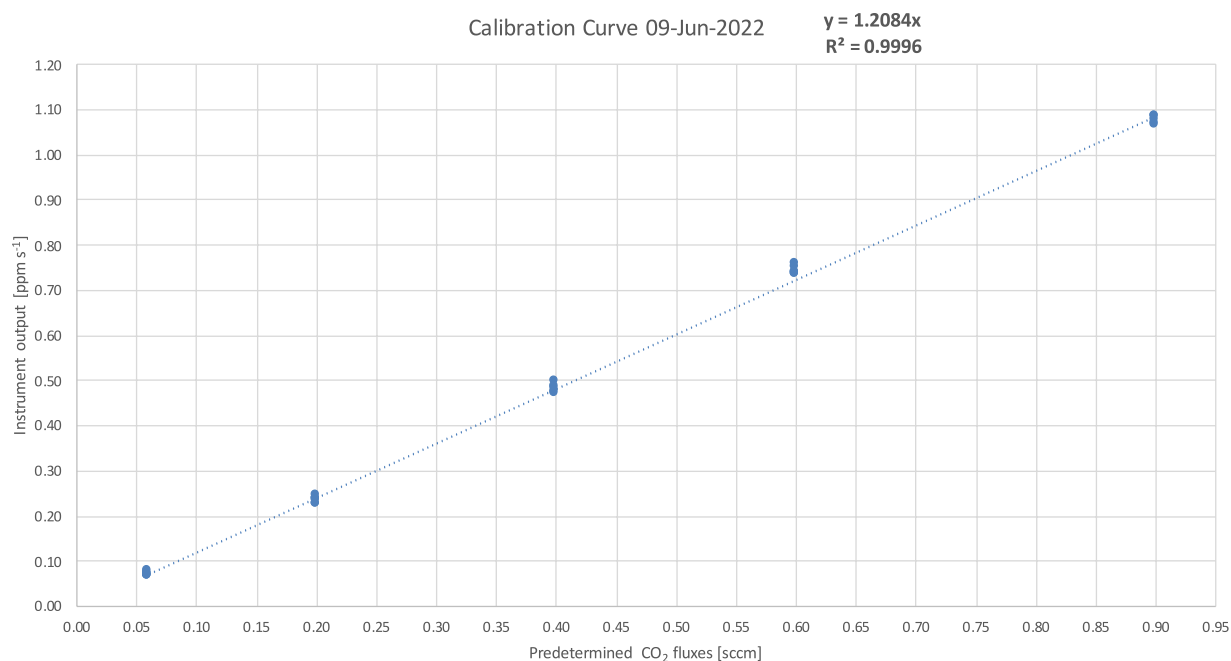


Fig. 8 Example of a calibration curve. Predetermined CO₂ flux values in [standard cc min⁻¹] are compared with the instrument outputs in [ppm s⁻¹].

2. Verifying the primary CO₂ span by measuring concentrations of 1.000 or 10.000 ppm CO₂.
3. Verifying the secondary CO₂ span by measuring near-ambient levels of CO₂.

To ensure high data quality, a data control process was conducted according to the outlined procedures. For each measurement campaign, the data were examined for anomalies or irregularities that could indicate

potential instrumental malfunctions, such as battery failure. The records corresponding to the identified critical issues were not removed from the dataset, rather the corresponding fields were reported as NA (Not Available). Furthermore, the correct application of formulas and calibration curves for each campaign to convert raw data from ppm s⁻¹ to μmolCO₂ m⁻²s⁻¹ was verified. This targeted approach to quality control aimed to preserve the rawness of the dataset, while addressing potential instrumental issues and thus ensuring correct data conversion.

Code Availability

No custom code was generated for this work.

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Competing interests

The authors declare no competing interests.

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