

1 **Evolving Particles in the 2022 Hunga Tonga–Hunga Ha’apai Volcano** 2 **Eruption Plume**

3 **Ralph A. Kahn^{1,2}, James A. Limbacher^{3,4}, Katherine T. Junghenn Noyes^{1,5}, Verity J.B.**
4 **Flower⁶, Lauren M. Zamora^{1,5}, and Kathleen F. McKee⁷**

5 ¹Senior Research Scientist Emeritus, NASA Goddard Space Flight Center, Greenbelt MD 20771.

6 ²Now at: The Laboratory for Atmospheric and Space Physics, University of Colorado Boulder,
7 Boulder CO 80303. ³I. M. Systems Group, Inc. (IMSG), Rockville MD 20850.

8 ⁴National Oceanic and Atmospheric Administration, College Park MD 20740

9 ⁵Earth System Science Interdisciplinary Center, University of Maryland, College Park MD,
10 20740

11 ⁶University of Stirling, Stirling, FK9 4LA Scotland

12 ⁷Vanderbilt University, Nashville TN 37235.

13
14 Corresponding author: Ralph A. Kahn (ralph.kahn@lasp.colorado.edu)

15 16 **Key Points:**

- 17 • Even for the optically thin, low-latitude Tonga plume, MISR retrieved aerosol plume-
18 height, associated wind vectors and particle properties
- 19 • Parts of near-tropopause and mid-stratosphere aerosol layers were observed downwind
20 over 10 days, nearly all spherical, non-light-absorbing
- 21 • The mid-stratosphere particles were smaller, but grew (ANG decreased by >20%)
22 between 1/17 and 1/23, consistent with model predictions
23

24 **Abstract**

25 The Multi-angle Imaging SpectroRadiometer (MISR) aboard NASA's Terra satellite observed
26 the Hunga Tonga—Hunga Ha'apai (HTHH) 15 January eruption plume on seven occasions
27 between 15 and 23 January 2022. From the MISR multi-angle, multi-spectral imagery we
28 retrieve aerosol plume height geometrically, along with plume-level motion vectors, and derive
29 radiometrically constraints on particle effective size, shape, and light-absorption properties. Parts
30 of two downwind aerosol layers were observed in different places and times, one concentrated in
31 the upper troposphere (11-18 km ASL), and a mid-stratosphere layer ~23 – 30+ km ASL. After
32 the initial day (1/15), the retrievals identified only spherical, non-light-absorbing particles,
33 typical of volcanic sulfate/water particles. The near-tropopause plume particles show constant,
34 medium-small (several tenths of a micron) effective size over four days. The mid-stratosphere
35 particles were consistently smaller, but retrieved effective particle size increased between 1/17
36 and 1/23, though they might have decreased slightly on 1/22. As a vast amount of water was also
37 injected into the stratosphere by this eruption, models predicted relatively rapid growth of sulfate
38 particles from the modest amounts of SO₂ gas injected by the eruption to high altitudes along
39 with the water (Zhu et al, 2022). MISR observations up to ten days after the eruption are
40 consistent with these model predictions. The possible decrease in stratospheric particle size after
41 initial growth was likely caused by evaporation, as the plume mixed with drier, ambient air.
42 Particles in the lower-elevation plume observed on 1/15 were larger than all the downwind
43 aerosols and contained significant non-spherical (likely ash) particles.

44

45 **Plain Language Summary**

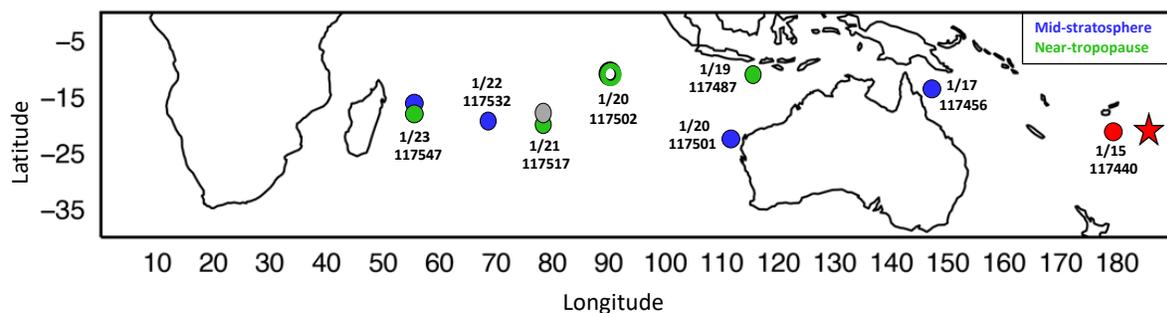
46 Volcanic eruptions often occur in remote locations, and typically present major hazards for
47 researchers and equipment. This creates opportunities for the frequent, global coverage satellite
48 remote sensing can make from the relative safety of space. Satellite instruments imaged the
49 massive 2022 Hunga Tonga—Hunga Ha'apai eruption plume, mapped the height of the initial
50 injections and distributions of water vapor, SO₂, and aerosol particles. The Multi-angle Imaging
51 SpectroRadiometer (MISR) instrument aboard NASA's Terra satellite offers some unique
52 capabilities in this respect, by observing the evolving particle properties in two aerosol layers,
53 one near-tropopause and one mid-stratosphere. For about a week following the initial
54 observations on 1/15, the only spherical, non-light-absorbing particles were retrieved by MISR.
55 The particles in the near-tropopause layer had relatively constant, several tenths of a micron
56 effective size. In the mid-stratosphere layer, the particles were systematically smaller, but
57 appeared to grow downwind. As a great deal of water vapor was also injected into the mid-
58 stratosphere, models predicted relatively rapid growth of sulfate particles from modest amounts
59 of SO₂ injected to high altitudes along with the water. The MISR results provide observational
60 support for these model predictions and some constraint on the associated rates, illustrating the
61 contribution satellite remote-sensing can make toward characterizing volcanic plume evolution.

62 **1 Introduction**

63 The Hunga Tonga—Hunga Ha'apai (HTHH) submarine volcano in the South Pacific
64 (20.6° S latitude; 175.4°W longitude) began erupting on 19 December 2021 (Matoza et al.,
65 2022); it produced a moderate eruption at ~03:20 UTC on 14 January; a major explosive event
66 began at ~04:14 UTC on 15 January that propelled considerable material well into the

67 stratosphere, with some reaching into the mesosphere (Amores et al., 2022; Carr et al., 2022;
 68 Klein 2022; Proud et al., 2022; Smart, 2022; Zuo et al., 2022). Wave perturbations were detected
 69 globally from ground stations and satellites, within the ocean, the land surface, and through the
 70 atmosphere up to the ionosphere (Matoza et al., 2022; Themens et al., 2022; Wright et al., 2022).
 71 Relatively little SO₂ and HCl were observed in the atmosphere compared to other major volcanic
 72 eruptions; for example, the 1991 Mt. Pinatubo eruption ejected ~40 times more SO₂ (e.g., Carn et
 73 al., 2022; Millan et al. 2022; Witze, 2022). However, HTHH introduced exceptionally large
 74 amounts of water vapor into the stratosphere, up to unprecedented elevation (up to ~58 km) and
 75 in quantities not previously observed during the satellite era (Millan et al. 2022; Silletto et al.,
 76 2022). Further, significantly more sulfate particles formed and propagated in the stratosphere
 77 than would be expected from the relatively modest SO₂ injection of this eruption (Legras et al.,
 78 2022; Taha et al., 2022).

79 The Multi-angle Imaging SpectroRadiometer (MISR) instrument aboard the NASA Earth
 80 Observing System's Terra satellite observed parts of the elevated aerosol plume on at least seven
 81 occasions between 15 and 23 January, under fair-to-good aerosol-retrieval conditions, as the
 82 plume traveled westward, traversing the Coral Sea, Australia, and the Indian Ocean (Figure 1).
 83 The MISR retrievals also provide some indication of interaction between the volcanic water and
 84 aerosol in the stratosphere as the plume evolved.



85

86 **Figure 1.** MISR Hunga Tonga—Hunga Ha’apai Volcano-plume-observation *Locations, Dates,* and Terra
 87 satellite *Orbit numbers*; the volcano is indicated with a red star. Observations of the mid-stratosphere layer, at
 88 ~23-30+ km, are shown in blue markers, those of the near-tropopause layer, at ~11-18 km, are shown in green.
 89 The red dot indicates where non-spherical particles were retrieved at 10-14 km within a day of the main
 90 eruption, most likely dominated by volcanic ash, and the gray dot shows where light-absorbing particles were
 91 retrieved, probably smoke originating from wildfires in Australia that might have mixed with the volcanic
 92 plume. All other retrievals reported spherical, non-light-absorbing particles, typical of volcanic sulfate/water
 93 particles. The open circle marks a downwind plume for which MISR plume heights were retrieved, but the
 94 AOD was too low for particle property retrievals.

95 The Terra satellite's near-polar orbit crosses the equator at approximately 10:30 AM local
 96 time. Each of the nine MISR cameras sweeps out a strip of imagery along the spacecraft track on
 97 the day side of Earth, successively at angles ranging from 70° in the forward direction, through
 98 nadir, to 70° toward the aft of the spacecraft, in each of four spectral bands centered at 446, 558,
 99 672, and 866 nm (Diner et al., 1998). It takes about seven minutes for all nine cameras to view a
 100 given location on Earth. The MISR swath is relatively narrow (~380 km), so sampling is
 101 sporadic, especially at low latitudes where the HTHH eruption occurred. Under good, but not
 102 necessarily ideal, retrieval conditions, qualitative particle microphysical properties can be

103 derived: three-to-five bins in retrieved, column-effective particle size (REPS), represented in the
104 retrieval results as the Ångström exponent (ANG) or qualitatively as “small,” “medium,” and
105 “large” particles between ~ 0.1 and ~ 2 microns effective radius. The exact boundaries of these
106 size bins depend on solar and viewing geometry, AOD, and particle properties, but differences in
107 REPS are much more robust, especially where the observing geometry is quite constant, as for
108 the cases in the current study (Kahn et al., 2001, Kahn and Gaitley, 2015). Under good retrieval
109 conditions, we can also retrieve two-to-four bins in column-effective particle light-absorption
110 (REPA), represented in the retrieval results as the single-scattering albedo at 558 nm wavelength
111 (SSA_{558}), along with spherical vs. randomly oriented non-spherical particles. For the current
112 study, we used the MISR Research Aerosol retrieval algorithm (RA), which offers 1.1 km pixel
113 resolution, the ability to co-register multi-angle views at plume elevation, advanced radiometric
114 calibration, and other refinements aimed at obtaining as much particle property information as
115 possible from the MISR radiances (Limbacher and Kahn, 2014; Limbacher et al., 2022). Due to
116 the stability of the instrument and the satellite, relative differences in REPS and REPA are much
117 more robust than the quantitative ANG and SSA values presented. Our conclusions in this study
118 rely primarily on relative differences in these quantities.

119 From a volcanological perspective, the particles below a few microns in size to which
120 MISR particle-property retrievals are sensitive represent the lower end of ash components
121 typically emitted during eruptions. However, the MISR-retrieved aerosol optical depth (AOD)
122 captures the extinction of all particles in the column, and the very fine ash and other aerosols
123 within the MISR size-retrieval-sensitivity range probably capture the tail of the full ash size
124 distribution. In themselves, these smaller components have the potential to produce far-reaching
125 effects, as they are the ones more likely to remain suspended in the atmosphere and be
126 transported long distances, impacting the radiative energy balance, air quality and health
127 downwind, as has been observed for other eruptions (e.g., Longo et al., 2010; Sahay et al., 2023;
128 Kluser et al., 2013).

129 MISR coverage is complemented by imagery from the broader-swath, single-view
130 MODerate resolution Imaging Spectroradiometer (MODIS) instrument, also aboard the Terra
131 satellite, and by layer-height curtains from space-based Cloud-Aerosol Lidar and Infrared
132 Pathfinder Satellite Observation (CALIPSO) lidar, where available. Where plume optical depth
133 is sufficient, 3-D aerosol plume-height maps can be derived geometrically from MISR multi-
134 angle imagery using the MISR Interactive eXplorer (MINX) software tool, with horizontal
135 resolution of 1.1 km and uncertainty in the vertical between 250 and 500 m (Nelson et al., 2013).
136 The MISR stereo-height retrievals have been validated against contemporaneous radar and lidar
137 observations (e.g., Naud et al., 2005; Marchand et al., 2007). MINX is currently configured to
138 retrieve plume altitude values from sea level to an elevation of 30 km. The plume-height
139 retrievals are keyed to the layer of maximum spatial contrast in the atmospheric column; this
140 generally identifies the optically thickest aerosol plume when multiple layers are present, unless
141 meteorological cloud or optically thick aerosol overlies other layers. As such, CALIPSO lidar
142 tends to identify thinner aerosols when they are present at higher elevations, whereas thermal
143 height-retrieval techniques tend to sample systematically lower than MISR/MINX (e.g., Flower
144 & Kahn, 2017). Aerosol plumes are rarely uniform, and MINX retrievals typically obtain a
145 range of elevations for a given plume; although the results are unlikely to capture the absolute
146 top or bottom of an aerosol layer, they often provide an indication of the plume vertical extent.
147 The associated wind vectors at plume elevation are also retrieved from MISR, based on the

148 actual movement of contrast elements in the aerosol plume itself over the seven minutes it takes
149 for all nine MISR cameras to image a given location. These wind vectors also allow for the
150 observed parallax to be corrected for plume proper motion, to produce “wind corrected” stereo
151 heights.

152 The constraints on particle size, shape, and light-absorption that can be derived from
153 MISR data depend upon the AOD and on the observed range of scattering angles (i.e., the angle
154 from the sun, down into the atmosphere and back up to the cameras). AOD at 558 nm (AOD_{558})
155 of at least 0.15 or 0.2, but also depending on surface brightness and variability, is required for
156 high-quality particle property retrievals (Kahn and Gaitley, 2015; Limbacher et al., 2022). The
157 HTHH plume data were acquired at nearly the solar equator ($\sim 20^\circ$ S latitude in January). This
158 means the range of scattering angles observed by MISR was about the smallest possible, which
159 limits the sensitivity of MISR to particle properties. Nevertheless, the available MISR data do
160 yield some constraints on particle properties and their evolution during the early days of plume
161 transport, as we present here.

162 In this paper, we summarize the MISR results, as a contribution to the overall picture of
163 this unique event. The instrument observed parts of the HTHH plumes with sufficient AOD to
164 obtain good retrieval results, in relatively cloud-free or broken cloud over-water areas, at least
165 seven times during the 10 days following the initial eruption. A near-source plume segment was
166 captured at an elevation of about 12 km ASL on 15 January (Figure 1, red dot), approximately 18
167 hours after the largest eruption began. On subsequent days, MISR observed parts of elevated
168 plumes toward the west, near the tropopause at $\sim 11 - 18$ km on 19, 21, and 23 January (Figure 1,
169 green dots), and well into the stratosphere at $\sim 23 - 30+$ km on 17, 20, 22, and 23 January (Figure
170 1, blue dots). For 20 January, the open green dot in Figure 1 shows the location of a plume
171 segment for which the plume height was retrieved, but the AOD was too low to derive particle
172 properties. On 21 January, in addition to the near-tropopause south plume segment, an area of
173 light-absorbing particles was retrieved (grey dot in Figure 1), most likely smoke from fires in
174 northern Australia. Both upper and lower layers were observed in the same region on 23 January.
175 In Section 2, we review the MISR plume height, wind vector, and particle property retrieval
176 results. We also use data from the CALIPSO space-based lidar (Winker et al., 2010) and from
177 several AERONET surface stations (Holben et al., 1998), each acquired in the general vicinity of
178 the MISR-observed plumes, to help assess the MISR plume height and particle property results,
179 respectively. In Section 3, we discuss the patterns and present some inferences about plume
180 evolution that we draw from the observations. Brief conclusions are given in Section 4.

181 **2 MISR Observations of the 2022 HTHH Eruption Plume**

182 In this section, we summarize each MISR observation of the HTHH plume between 15
183 and 23 January 2022, providing our interpretation of the MINX and RA results, including
184 comparisons with the closest AERONET and CALIOP data, where available. We focus on ANG
185 rather than the individual retrieved components, as qualitative differences in ANG are more
186 robust than the components and proportions that the algorithm identifies as providing a match to

187 the observed top-of-atmosphere reflectances. The candidate particle optical models used for the
 188 RA retrievals presented here are given in Table S1 in Supplemental Material.

189 We took several additional steps to maximize confidence in the REPS and REPA results.
 190 The cost function used to determine acceptable RA retrieval results is given as:

191

$$\text{Cost} = \frac{\sum_{\lambda} \sum_c \left(\frac{\sqrt{w_{\lambda,c}} [BRF_{\lambda,c}^{\text{TOA}} - BRF_{\lambda,c}^{\text{mod}}]}{\text{Unc}_{\lambda,c}} \right)^2}{\sum_{\lambda} \sum_c w_{\lambda,c}} \quad (1)$$

192

193 where BRF^{TOA} are the MISR-observed bi-directional reflectance factors as a function of
 194 wavelength (λ) and camera (c), BRF^{mod} are the model-simulated bi-directional reflectance factors
 195 for a given aerosol mixture, and wavelength-dependent optical depth, surface albedo, and
 196 atmospheric backscatter fraction, Unc are estimated, channel-specific radiance uncertainties, and
 197 w are channel-specific weights that account for topographic shadowing, any glint contamination
 198 over water, and missing data (details are given in Limbacher & Kahn, 2019 and Limbacher et al.,
 199 2022). We examined the dependence of ANG and SSA results on the value of the cost function;
 200 for cost function thresholds of 2 and 1 the results are similar in all cases, and the higher threshold
 201 provides greater coverage. So, we applied a cost function upper bound of 2 to all the results
 202 presented in this paper except for the optically very thick, near-volcano 15 January case.

203 At low AOD, retrieval sensitivity to particle properties diminishes (e.g., Kahn and
 204 Gaitley, 2015). To investigate the relationship between AOD and retrieved particle properties,
 205 we plotted the dependence of ANG and SSA on the retrieved AOD. These plots are included as
 206 Figures S1 and S2 in the Supplemental Material for the near-tropopause and mid-stratosphere
 207 cases, respectively. Retrieved SSA is essentially independent of AOD in all cases after 15
 208 January, whereas ANG tends to increase with increasing AOD_{558} . The ANG variation appears to
 209 be real, associated with geographical variation in plume properties shown in subsequent figures.
 210 We set a lower bound of 0.15 on AOD_{558} for reporting ANG and SSA statistics to exclude areas
 211 where retrieval uncertainty is highest (though this affects only the 21 January case), and an upper
 212 bound of 8.0 to help reduce cloud contamination.

213 Applying these criteria, we examine the seven MISR observations of the HTHH plume:

214

215 **Table 1.** Overview of MISR HTHH Plume Observations

Observation date*	Observation time [UTC]	Approx. Distance from Source [km]	Retrieved Height [km ASL]	Retrieved Zonal winds [†] [m/s]	Median AOD_{558}	Median ANG
15-Jan	~22:05	340	10 to 14	-10	~3	<~0.4
17-Jan	~00:25	2700	27 to 30+	28	~0.7	~1.4
19-Jan	~2:30	6500	11 to 16	24	~1.1	~0.86
20-Jan	~2:38	7400	24 to 29	32	~0.5	~1.2

21-Jan	~5:00	10750	23 to 26	24	~0.5	~0.94
22-Jan	~5:43	11800	23 to 26	25	~0.4	~1.3
23-Jan	~6:26	12800	15 to 17	12	~0.4	~1.0

216 *19-Jan main retrieval region; 20-Jan Orbit 117501; 21-Jan South Plume

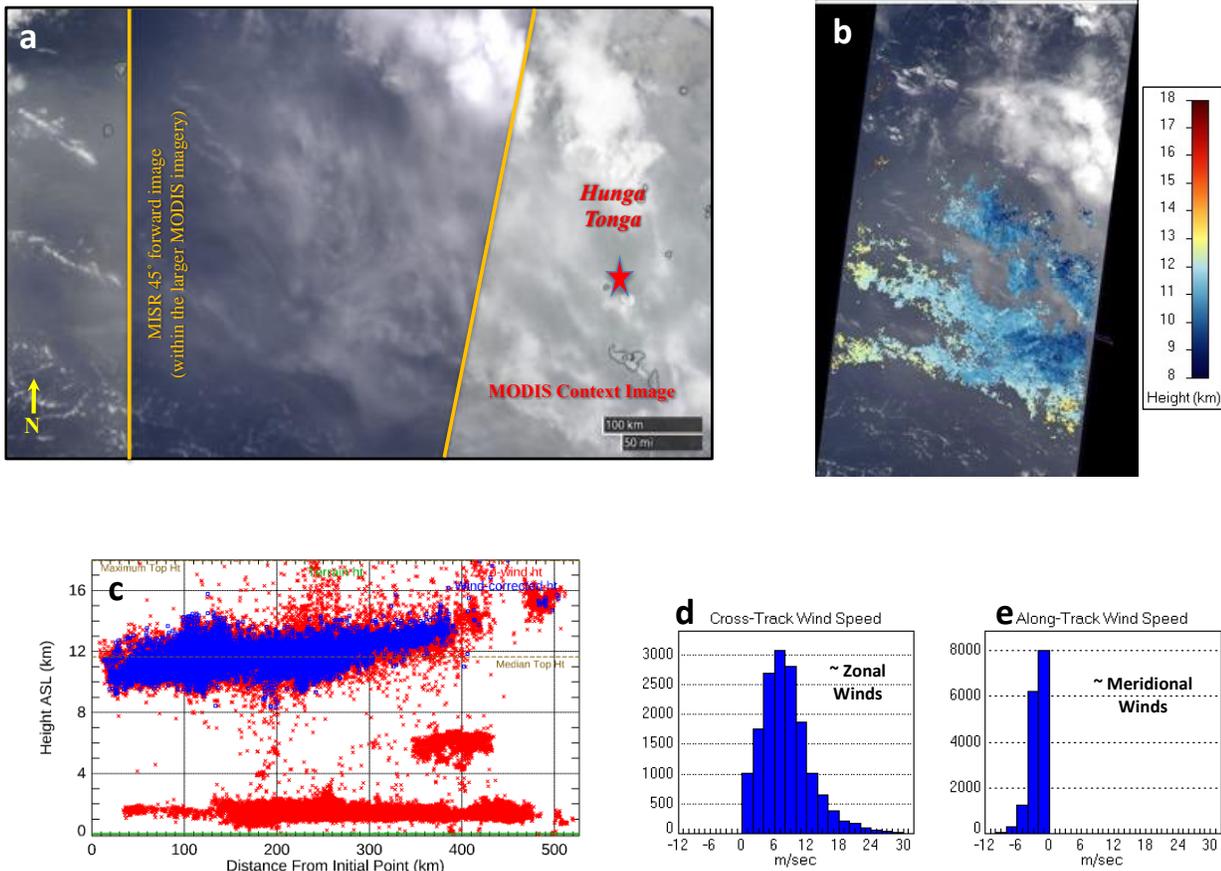
217 †Zonal winds are approximate westward median values

218

219 *15 January 2022, ~22:05 UTC.* MISR first imaged the HTHH plume at this time, capturing a
 220 swath centered approximately 340 km west of the volcano itself. Figure 2 shows the MISR-
 221 retrieved height for this plume segment at 10–14 km ASL. Careful cloud-screening in the MISR
 222 RA limits the number of particle property retrievals in the MISR region over which MINX
 223 plume heights shown in Figure 2 were derived. High-quality RA retrievals were obtained only on
 224 the west side of the study region shown in Figure 3 for this day. AOD_{558} is ~3, the column-
 225 effective ANG is $< \sim 0.4$, indicating a significant contribution from medium-large particles (i.e.,
 226 $> \sim$ a micron in radius), the non-spherical AOD fraction is ~0.16, and the SSA_{558} is ~0.99, all
 227 suggesting the presence of some larger, non-spherical volcanic ash or ice particles (see also
 228 Table S1). Depolarization and color-ratio observations from CALIPSO also suggest the
 229 occurrence of non-spherical ash or ice at elevations up to 40 km on this day (Sellitto et al., 2022).
 230 A comprehensive analysis of the 15 January 2022 plume dispersion during the first ~18 hours,

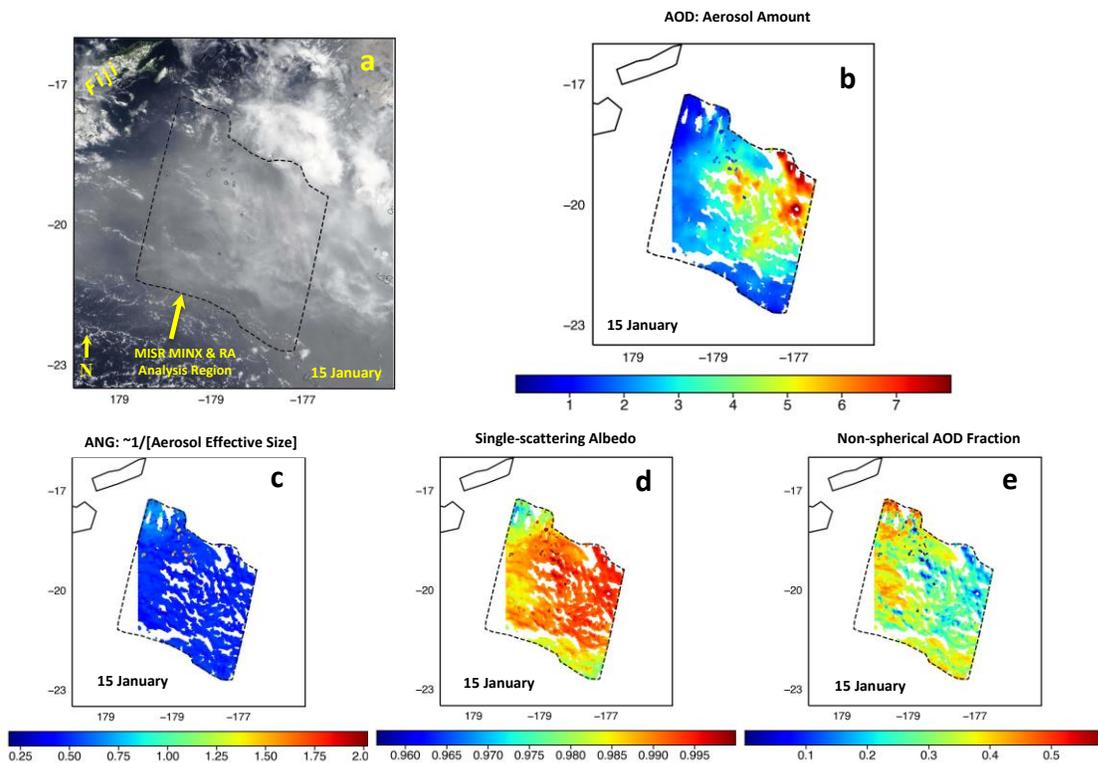
231 including these MISR observations and measurements from other sources, is part of continuing
 232 work.

233 On all subsequent days, discussed below, MISR retrieved only spherical, non-light-
 234 absorbing particles in downwind plume segments, at elevations from near the tropopause up to
 235 the MINX elevation retrieval limit of 30 km ASL.



236

237 **Figure 2.** MISR MINX Plume-height and Wind Vector retrieval results for 15 January 2022 at ~22:05 UTC,
 238 Terra Orbit 117440. (a) MODIS context image, with MISR 45° forward view superposed (between the orange
 239 lines). The volcano location is indicated by a red star. (b) MISR MINX plume-height map, superposed on
 240 MISR nadir-view image. (c) MISR-retrieved plume-height profile, starting point on the east side of the MISR
 241 field-of-view. (d) Cross-track (roughly west-east) and (e) along-track (roughly south-north) MISR-retrieved
 242 wind vector components, assessed over all plume points in the digitized region. (Although this plume segment
 243 was observed by MISR to the west of the volcano, the retrieved positive zonal wind indicates eastward motion.
 244 The zonal winds varied in elevation and time on this day, and an eastward component at the location and time
 245 of the MISR observation is supported by National Center for Physical Acoustics reanalysis (NCPA,
 246 <https://g2s.ncpa.olemiss.edu/>). For all other cases in this study, the MISR-retrieved zonal wind is toward the
 247 west.) In Panel (c), Red = zero-wind height; Blue = wind-corrected height; Green = surface elevation. No
 248 wind-corrected heights were derived for the two lower layers at ~6 km and ~11 km elevation in Panel c, as the
 249 origin of digitization was selected for the more prominent layer at ~12 km.

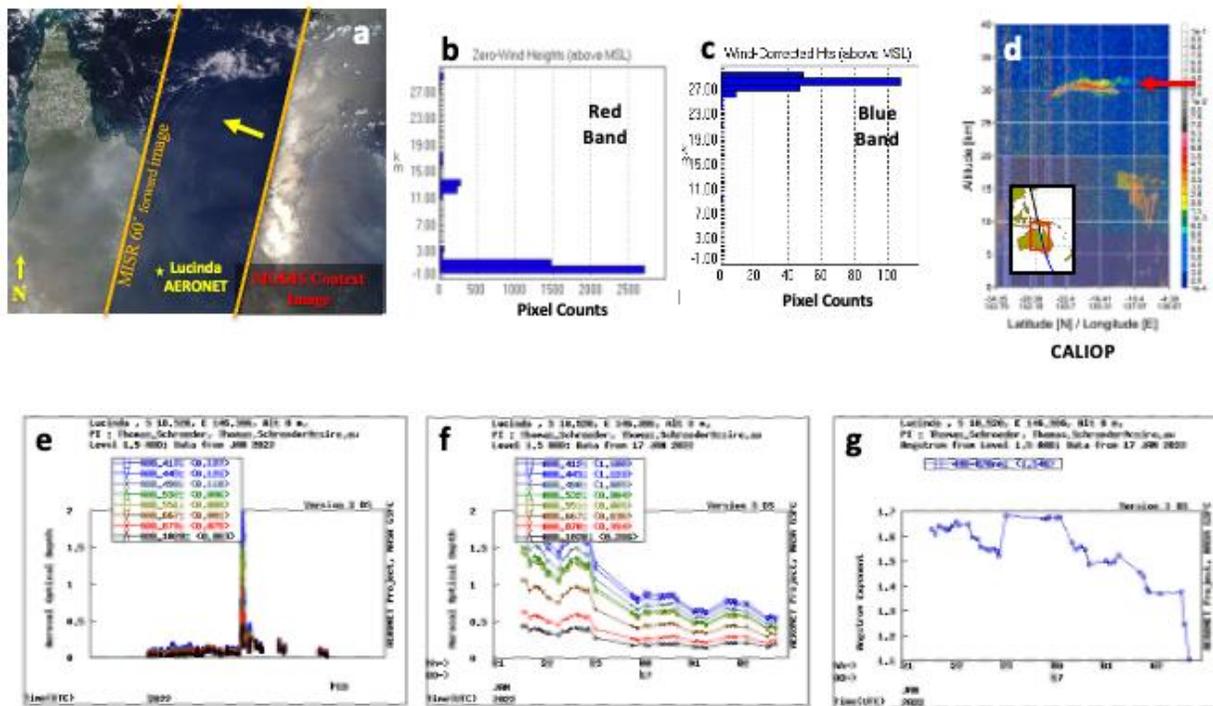


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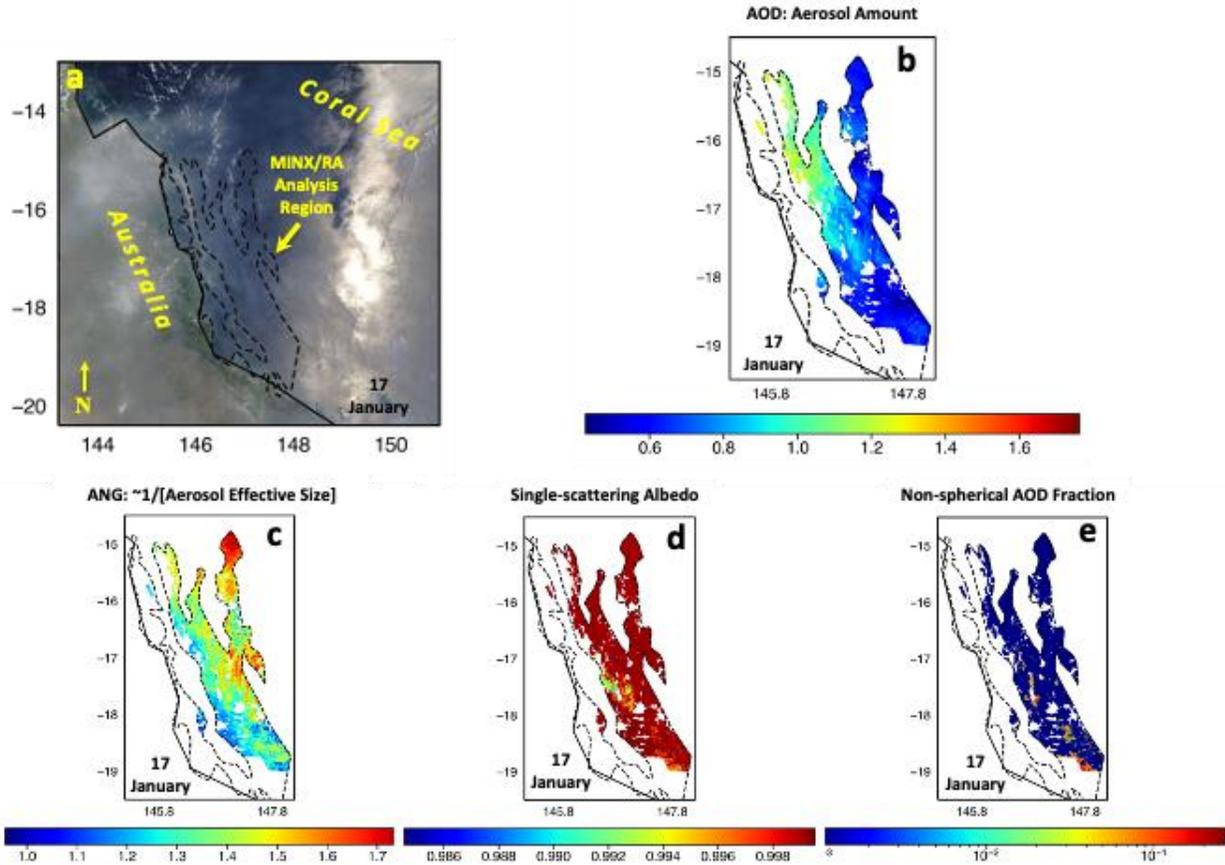
251 **Figure 3.** MISR Research Algorithm particle property retrieval results for 15 January 2022, Terra Orbit
 252 117440. (a) MISR context image with study region outlined; the relationship of this region to the HTHH
 253 volcano is shown in Figure 2a. (b) AOD₅₅₈ in the less cloudy parts of the study region; (c) ANG assessed
 254 between 446 and 866 nm; (d) SSA₅₅₈; (e) Non-spherical AOD₅₅₈ fraction. No cost-function filter was applied
 255 in this one case, but to improve confidence in the particle property retrieval results, ANG and SSA were
 256 assessed over the AOD range $0.15 < \text{AOD}_{558} < 8.0$. Digital data for the MISR RA results are given in Table S1.

257 *17 January 2022, ~00:25 UTC.* On this day, MISR captured a part of the HTHH plume off the
 258 northeast coast of Australia, ~2,700 km west of the volcano, and ~450 km north of the Lucinda
 259 AERONET site. Plume heights reached from 27 and 30+ km ASL (Figure 4). (As 30 km is the
 260 highest elevation allowed by the MINX algorithm, some contrast elements might have occurred
 261 above the MINX-reported peak altitude.) AOD₅₅₈ had dropped from the previous MISR plume
 262 observation to ~0.7, with ANG ~1.4 (Figure 5 and Table S1). Such small particles provided
 263 enough opacity in the MISR blue band (at 446 nm) for contrast features to allow a plume-height
 264 retrieval for the 30+ km layer (Figure 4c). However, the elevated plume particles were
 265 apparently too small and optically thin in the MISR red band (at 672 nm) to produce sufficient
 266 contrast, and only a near-surface layer and a weaker layer around 12 km were retrieved in this
 267 channel (Figures 4b). CALIOP lidar confirms the presence of the ~30 km layer in this general
 268 area (Figure 4d). The AERONET station at Lucinda obtained contemporaneous measurements

269 of AOD₅₅₀ ~0.7 and ANG ~1.55 (Figures 4f and 4g), confirming to the extent possible the
 270 presence of an aerosol layer similar to that retrieved by MISR.



271
 272 **Figure 4.** MISR MINX Plume-height retrieval results for 17 January 2022, Terra Orbit 117456,
 273 ~00:25 UTC, acquired north of Queensland, Australia, along with Lucinda AERONET ground-
 274 station retrievals. (a) MODIS context image, with MISR 60° forward view superposed (between
 275 the orange lines). The MISR retrieval region is indicated by the yellow arrow. The location of the
 276 Lucinda AERONET station on the northeast Australia coast, marked with a yellow asterisk
 277 within the MISR swath, ~450 km south of the over-ocean region where the MISR retrievals were
 278 performed, is still apparently within the plume. (For scale, the MISR swath is roughly 380 km
 279 wide.) (b) Red-band MISR-MINX plume-height histogram, showing a weak layer near the
 280 tropopause at 12-14 km and another near-surface. (c) Blue-band MISR plume-height histogram,
 281 showing a layer at 27-30+ km. (d) CALIOP Level 1, version 4.51, total attenuated backscatter
 282 ($\text{km}^{-1} \text{sr}^{-1}$) at 532 nm, acquired at ~05:28 UTC on 17 January, roughly 5 hours after and 800 km
 283 downwind of the MISR observation at the nearest point. (The inset gives the orbit location, and
 284 the red arrow highlights the likely plume layer.) (e) Lucinda AERONET AOD for 24 January
 285 2022, showing the AOD peak on 17 January. (f) Lucinda AERONET AOD for 17 January 2022.
 286 (g) Lucinda AERONET Ångström Exponent for 17 January 2022. (AERONET data were
 287 provided by Thomas Schroeder and the AERONET team.)



288

289 **Figure 5.** MISR RA particle property retrieval results, same as Figure 3, but for 17 January 2022, Terra Orbit
 290 117456. Note that the scale for panel (e) is logarithmic here to show differences, given the range of in values.

291 *19 January 2022, ~2:30 UTC.* MISR observed a plume segment southwest of Java (Figure S3 in
 292 Supplemental Material), yielding a plume height of 11-16 km and MINX zonal winds at plume
 293 elevation in excess of 24 m/s toward the west. Neither CALIPSO nor ICESat-2 have
 294 measurements sufficiently proximal to the MISR observations in space and time to provide
 295 validation of the MISR height retrievals. At these low latitudes, the MISR layer would be below,
 296 but in the vicinity of, the tropopause elevation. In the least cloud-affected part of the field-of-
 297 view, the MISR RA retrieved AOD₅₅₈ ~1.1, ANG ~0.8, with spherical, non-light-absorbing
 298 particles dominating (Figure S4 and Table S1).

299 *20 January 2022, ~2:38 UTC.* On 20 January MISR observed part of the HTHH plume along
 300 the west coast of Australia, about 7,400 km west of the volcano. Here the plume was elevated
 301 between 24 and 29 km ASL, with zonal winds of ~30 m/s (Figure S5). CALIOP identified a
 302 plume at ~25 – 27 km ASL about 7 hours earlier and 400 km downwind of the MISR
 303 observation. MISR median AOD₅₅₈ was about 0.5, most ANG values ~ 1.1, and the particles
 304 were again spherical and non-light-absorbing (Figure S6 and Table S1). For comparison, the
 305 Learmonth AERONET, about 350 km northeast of the MISR study region along the Australia
 306 west coast, obtained AOD₅₀₀ and ANG at the closest observations to MISR overpass time of 0.7
 307 and 1.05, respectively (Figures S6g and S6h). The somewhat larger AOD at the AERONET site
 308 could be due to heterogeneity in the aerosol plume after a 5-day transit, with possible
 309 contributions from continental aerosol from Australian sources. There was a second MISR

310 overpass of a part of the likely volcanic plume ~2,700 km further west, at ~4:14 UTC. The
311 MISR-MINX retrieval obtained a plume height of 13-17 km ASL, with the hint of another layer
312 at 25-28 km ASL (Figure S7); however, the AOD was too low to derive meaningful particle
313 microphysical properties from the data.

314 *21 January 2022, ~5:00 UTC.* The plume observed by MISR on this day was in the central
315 Indian Ocean, about 10,750 km from the source. MISR retrieved a consistent plume elevation of
316 23-26 km ASL (Figure S8). (ICESat-2 indicated several lower-level aerosol layers between 8 and
317 12 km, but no retrievals are currently available that sample above 14 km.) The MISR study
318 region on this day contained two plume segments, north and south, that showed distinctly
319 different particle microphysical properties. As with most of the other downwind plume parts
320 captured by MISR, the South segment on 21 January contained small-medium, spherical, non-
321 light absorbing particles, with $AOD_{558} \sim 0.5$, $ANG \sim 0.95-1.0$, and $SSA_{558} \sim 1.0$ (Figure S9 and
322 Table S1).

323 However, for the North segment, AOD_{558} is ~ 0.9 , ANG is $\sim 1.4-1.5$, and SSA_{558} is
324 between 0.94 and 0.96 (Figure S10 and Table S1). The low SSA_{558} and smaller effective particle
325 size suggest the North segment might contain smoke. This seems likely, as major fires were
326 burning in northeast and northwest Australia, producing extensive smoke plumes visible in
327 MODIS imagery from 16 to 18 January (<https://worldview.earthdata.nasa.gov>; last accessed 01
328 May 2023). Trajectory analysis from NOAA's HySPLIT model (Rolph et al., 2017) indicates
329 that transport from across north Australia to the location of the MISR plume observation on 21
330 January ($\sim 21.4^\circ S$, $77.3^\circ E$) occurred in 3-4 days at elevations from 18 to over 27 km. This result
331 also represents a demonstration of our ability to derive distinct aerosol type differences from the
332 MISR multi-angle, multi-spectral data.

333 *22 January 2022, ~5:43 UTC.* MISR observed a plume segment at 23-26 km ASL on 22
334 January, about 11,800 km from the source (Figure S11). Yet, the MISR RA derived a mean
335 AOD_{558} around 0.4, within the range for which particle properties can be retrieved over dark
336 water (Figure S12 and Table S1). The particles are all of small size ($ANG \sim 1.3$) and are still
337 spherical and non-light absorbing, as for all previous downwind retrievals associated with the
338 volcano plume.

339 *23 January 2022, ~6:26 UTC.* This was the final occasion on which MISR captured a segment
340 of the HTHH eruption plume having sufficient definition to report plume elevations with some
341 confidence. In this case, we derive particle microphysical properties qualitatively in two distinct
342 layers. By the time of these observations, the plume particles had reached the vicinity of Réunion
343 Island, $\sim 12,800$ km west of the volcano and ~ 720 km east of Madagascar. The lower layer was
344 concentrated 15-17 km ASL, whereas the thinner upper layer concentrated at $\sim 24-26$ km above
345 the ocean surface, extending over the lower layer as well as to the northwest and southeast of the
346 lower layer (Figure S13b, c). The MISR retrieval indicates westward motion at $\sim 8-12$ m/s
347 (Figure S13e, f). For the "mixed-height" layer, dominated by the lower, optically thicker layer,
348 total-column AOD_{558} was retrieved at ~ 0.4 and $ANG \sim 1.0$ (Figure S14 and Table S1). For the
349 elevated layer alone, AOD_{558} was $< \sim 0.3$ and $ANG \sim 1.1$. The AERONET site at Réunion, just

350 north of the MISR retrieval region, reported AOD500 ~0.4 and ANG ~0.63 at the 6:26 UTC
351 MISR overpass time, i.e., larger particles than those observed by MISR (Figure S13).

352 There was also a stratospheric balloon flight deployed from La Réunion island itself
353 (21.1° S; 55.3° E) on this day (20:04 – 21:35 UTC), carrying an optical counter with sensitivity
354 to particle sizes from 0.2 μm to ~30 μm diameter in 19 bins (Kloss et al., 2022). They also had
355 a ground-based backscatter lidar. The balloon identified plume layers at 22.6 and 24.9 km ASL,
356 the upper layer corresponding to the upper layer identified by MISR (Fig. S13c, d). The lower
357 layer identified by MISR is about 6 km below the lower layer reported by the *in situ*
358 measurements. (There are also discrepancies in layer elevations between the balloon and the
359 ground-based lidar, likely due to differences in spatial sampling combined with layer
360 heterogeneity.) Consistent with the MISR results, the optical counter found spherical, non-light
361 absorbing particles with radii $<0.5 \mu\text{m}$ in the upper layer. In their 22.6 km layer, they report
362 some light-absorbing particles, which could be transported smoke, as we observed with MISR on
363 21 January.

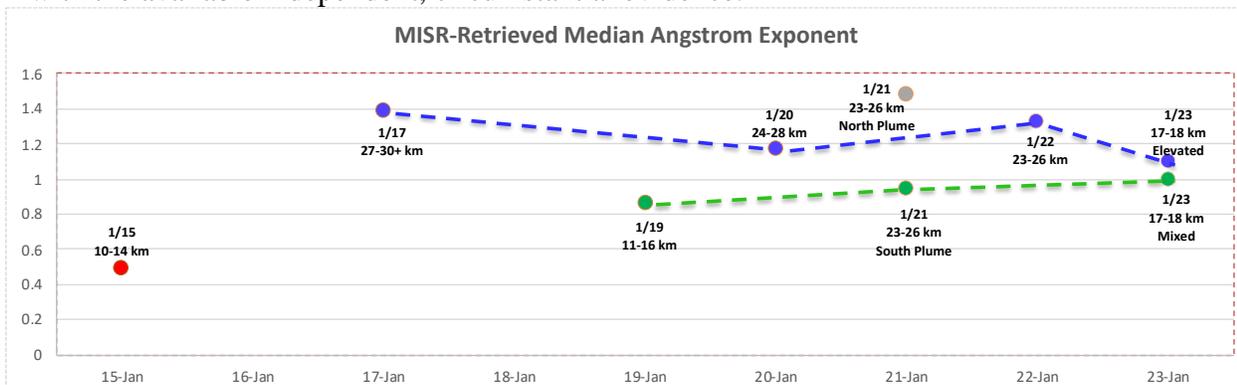
364 **3 Discussion – MISR Constraints on the Evolution of the HTHH Aerosol Plume**

365 The altitude reached by volcanic emissions depends on ejection force, plume buoyancy,
366 atmospheric stability profile, wind shear, and particle properties. Generally, ash particles are
367 likely to settle more rapidly than smaller, lighter, aerosol components, especially those that tend
368 to form after eruption from water vapor and other gases, such as typical sulfate/water particles in
369 volcanic plumes. On 15 January, MISR retrieved particles containing a significant fraction of
370 non-spherical components, with ANG <0.4 , indicating medium-large particles (Table S1), most
371 likely dominated by volcanic ash. A complimentary study by *Sellitto et al. (2022)* concluded that
372 ash was deposited rapidly between 19:00 UTC on 15 January and 02:00 UTC on 16 January,
373 based on frequent observations made with the geostationary Himawari 8 satellite. The MISR
374 observation, obtained at 22:05 UTC on 15 January, falls within this window. As such, the MISR
375 observation was likely acquired during a period of rapid ash deposition. MISR ash detection on
376 this day is further supported by the lack of ash markers identified in subsequent MISR retrievals;
377 between 17 and 23 January MISR retrievals consistently found spherical, non-light-absorbing
378 aerosols, more likely composed of sulfate/water than ash. (The other exception occurred on 21
379 January, when in addition to the volcanic plume, MISR observed what was likely transported
380 smoke (Figure S10).) We note that *Colombier et al. (2023)* posit non-spherical sea salt particles
381 might be part of the 15 January atmospheric aerosol load, given the submarine nature of the
382 eruption and the presence of sea salt in surface samples acquired on nearby islands. With MISR
383 we cannot distinguish such particles from ash, so our data have nothing to contribute about this
384 possibility.

385
386 Other satellite observations identified an SO₂ gas plume ejected during the HTHH
387 eruption and tracked the resulting sulfate particles to altitudes exceeding 45 km (*Sellitto et al.,*
388 *2022; Taha et al., 2022*), possibly up to where the plume achieved neutral buoyancy and/or
389 where ambient conditions, such as the anomalously high concentration of water vapor, favored
390 rapid sulfate formation through oxidation and hydration processes. The apparent vertical
391 separation of sulfate/water and ash plumes has also been observed in other volcanic eruptions,
392 such as Grimsvotn in 2011 (*Flower & Kahn, 2020*). Preferential settling of ash likely contributed
393 to the lack of larger, non-spherical particles observed further downwind from the volcano after

394 15 January. The observation of localized ash-fall around the volcano immediately following the
 395 eruption (e.g., *Thompson, 2022*) also supports the rapid-ash-deposition hypothesis.

396
 397 Two aerosol layers were observed in the downwind MISR plume detections, one near-
 398 tropopause and one mid-stratosphere. For the near-tropopause layer, AOD₅₅₈ exceeded unity on
 399 19 January, diminishing to ~0.47 on 21 January (South plume) ~0.4 on 22 January, and 0.26 by
 400 23 January in the MISR retrievals (Table S1). These retrievals indicate spherical, non-light-
 401 absorbing particles, with median ANG between 0.85 and about 1.0, i.e., small-medium particles.
 402 possibly decreasing very gradually in effective size (Figure 6). With the sparse MISR sampling
 403 of these extended aerosol layers, plume heterogeneity is likely, so the values reported should be
 404 considered qualitatively representative of overall plume properties during transport, with higher
 405 confidence in the relative changes than the absolute values. The North plume segment on 21
 406 January has distinctly different microphysical properties than the adjacent South plume segment
 407 observed at the same elevation: significantly higher median AOD₅₅₈ (~0.9), lower SSA₅₅₈
 408 (~0.96), and larger ANG (~1.5). The North plume properties are typical of transported wildfire
 409 smoke, and model trajectories from burning regions in parts of north Australia support this
 410 suggestion (see Section 2 above). These results also contribute to validating the qualitative MISR
 411 particle microphysical property retrievals, as they show distinctions that are at least consistent
 412 with the available independent, circumstantial evidence.



413 **Figure 6.** MISR-retrieved Ångström exponent (ANG, inversely proportional to effective particle size) for the
 414 MISR observations of the near-source HTHH plume on 15 January (red dot), and for the near-tropopause
 415 (green dots) and stratosphere (blue dots and gray dot) plumes downwind. Observation dates and plume heights
 416 are given in the annotations.
 417

418
 419 AOD₅₅₈ in the mid-stratosphere layer diminished systematically from about 0.7 to ~0.26
 420 between 17 and 23 January (Table S1). Essentially spherical, non-light-absorbing particles were
 421 retrieved in all cases for this layer. Of particular interest is the effective particle size, which is
 422 inversely proportional to the MISR-retrieved ANG plotted for the available cases in Figure 6.
 423 For the plume segments in the mid-stratosphere layer (blue dots), the effective particle size was
 424 always smaller than that for the near-tropopause layer (green dots), and the retrieved effective
 425 particle size (REPS) actually increased between 17 and 20 January, then appears to have
 426 decreased by 22 January and possibly increased again in the lowest-AOD observation on 23
 427 January. From the available snapshots, particle evolution can only be inferred; if these
 428 observations capture the evolution of particles in the layer near 30 km, either coagulation of
 429 particles in the layer or the condensation of water might be responsible for the observed particle
 430 growth. The layer mid-visible optical depth was ~0.4 on 22 January (Table S1) and the layer

431 vertical extent was at least 5 km (Figure S11c), so particle concentrations were fairly low, which
432 does not favour particle collision and coagulation. However, an unprecedented amount of water
433 vapor was present at these elevations; ordinarily, the stratosphere at ~30 km is very dry, whereas
434 HTHH was a “wet” eruption, and Microwave Limb Sounder (MLS) data show that significant
435 water was injected to these elevations (*Millan et al.*, 2022). This supports an explanation for the
436 apparent particle growth as the result of particle hydration, either by adsorption or by ice
437 deposition. Ice deposition on volcanic particles is a well-known phenomenon in wet (i.e.,
438 phreatomagmatic) eruption plumes (e.g., Van Eaton et al., 2012), an example of the classic
439 juxtaposition of fire and ice.

440 **4 Conclusions**

441 As the MISR swath-width is relatively narrow, inferences drawn about HTHH plume
442 evolution from the snapshots of this low-latitude eruption presented here must be considered in
443 the context of broader-swath observation provided by other instruments, as cited above. Yet,
444 MISR does contribute unique information, particularly about changes in height-resolved particle
445 properties during the week following the eruption. The MISR-retrieved spectral AOD results
446 generally align with those obtained by ground-based sun photometers, which provide some
447 validation for the MISR aerosol amount (AOD) and particle size (ANG) retrievals. Similarly,
448 three cases where the CALIPSO lidar observations of plume height were sufficiently close in
449 space and time to the MISR observations support the MISR-MINX plume-height results.

450
451 MISR captured two distinct layers of the HTHH eruption plume downwind, one near-
452 tropopause and one mid-stratosphere. On 15 January, a plume rich in medium-large, non-
453 spherical particles (likely volcanic ash) was observed a few hundred kilometers to the west of the
454 volcano. Thereafter, only small or small-medium, spherical, non-light-absorbing particles,
455 typical of volcanic sulfate/water, were retrieved by MISR. By 23 January, the sulfate-like HTHH
456 aerosol plume had remained suspended for more than a week, had travelled at least a third of the
457 way around the globe, and was still sufficiently optically thick to support MISR plume-height
458 and particle-property retrievals, though with somewhat lower confidence than the higher-AOD
459 cases.

460
461 Particles in the mid-stratosphere layer (blue dots in Figures 1 and 6) were systematically
462 smaller than the near-tropopause particles (green dots). The mid-stratosphere particles apparently
463 grew in size between 17 and 20 January, likely due to hydration or water condensation, then
464 seem to have fluctuated in size between 22 and 23 January, possibly due to some combination of
465 measurements sampling different parts of a heterogeneous plume and evaporation as the plume
466 mixed with drier air and dissipated downwind. For the near-tropopause layer, particle size
467 remained relatively constant or very gradually diminished between 19 and 23 January. Due to the
468 stability of the instrument and the satellite, these relative size differences, retrieved from MISR,
469 are much more robust than absolute values. As such, our main conclusions rely on relative
470 differences in particle microphysical properties and show consistency with model predictions
471 (e.g., Zhu et al., 2022).

472
473 We note also that after we originally submitted the current paper, several preliminary
474 studies appeared relevant to key aspects of our conclusions. Specifically, Biochu et al. (2024)
475 combined HIMAWARI-8 geostationary thermal infrared imagery with CALIPSO lidar to

476 identify ash on 15 January, and subsequently smaller, but initially rapidly growing, stratospheric
477 aerosol with estimated 0.3-0.5 μm radius, which they interpret as sulfate and track for up to 1.4
478 years thereafter. They support their satellite-data interpretation with measurements from multiple
479 low-latitude AERONET stations, including the Lucinda and Learmonth stations that we use for
480 validation in the current study. Gupta et al. (2024) used Stratospheric Aerosol and Gas
481 Experiment-III (SAGE-III) observations to characterize stratospheric aerosol amount and
482 extinction associated with the HTHH eruption; they highlight the importance of knowing the
483 sulfate aerosol amount and properties, showing that although water vapor added to the
484 stratosphere would produce surface warming, this would be offset by a net cooling due to
485 stratospheric ozone loss and sulfate aerosol production.

486
487 The MISR results presented here demonstrate our space-based multi-angle, multi-spectral
488 aerosol retrieval capabilities, and represent the MISR contribution to ongoing investigations by
489 the wider community aimed at deriving as much as possible about HTHH plume-particle
490 evolution from satellite remote-sensing data and modelling.

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498 archiving the Level 1 MISR data used in this study.

499

500 **Open Research**

501 MISR data are freely available for download from NASA's MISR data repository
502 (<https://10dup05.larc.nasa.gov/MISR/cgi-bin/MISR/main.cgi>), maintained by the NASA Langley
503 Research Center (LaRC) Atmospheric Science Data Center (ASDC). The MISR INteractive
504 eXplorer (MINX) program, for determining plume altitude and associated motion vectors, is a
505 stand-alone software package developed at the NASA Jet Propulsion Laboratory (JPL) and
506 distributed through Github (<https://github.com/nasa/MINX/releases>). MODIS true color imagery
507 are accessed through the NASA Worldview application (<https://worldview.earthdata.nasa.gov>),
508 part of the NASA Earth Observing System Data and Information System (EOSDIS). The
509 AERONET and CALIPSO data are freely available from their respective archives at
510 <https://aeronet.gsfc.nasa.gov> and <https://subset.larc.nasa.gov/calipso/>. MISR RA and MINX
511 results generated in this work will be accessed through the NASA Langley ASDC, where the
512 MISR Standard Products are also archived (<https://eosweb.larc.nasa.gov/project/MISR>). All the
513 results of the study are presented in the paper itself and supplemental material, including the key
514 data statistics associated with this study.

515

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678

679 **Figure Captions**

680

681 **Figure 1.** MISR Hunga Tonga—Hunga Ha’apai Volcano-plume-observation *Locations, Dates,*
 682 *and Terra satellite Orbit numbers;* the volcano is indicated with a red star. Observations of the
 683 mid-stratosphere layer, at ~23-30+ km, are shown in blue markers, those of the near-tropopause
 684 layer, at ~11-18 km, are shown in green. The red dot shows where non-spherical particles were
 685 retrieved at 10-14 km within a day of the main eruption, most likely dominated by volcanic ash,
 686 and the gray dot shows where light-absorbing particles were retrieved, probably smoke
 687 originating from wildfires in Australia that might have mixed with the volcanic plume. All other
 688 retrievals reported spherical, non-light-absorbing particles, typical of volcanic sulfate/water
 689 particles. The open circle marks a downwind plume for which MISR plume heights were
 690 retrieved, but the AOD was too low for particle property retrievals.

691

692 **Figure 2.** MISR MINX Plume-height and Wind Vector retrieval results for 15 January 2022 at
 693 ~22:05 UTC, Terra Orbit 117440. (a) MODIS context image, with MISR 45° forward view
 694 superposed (between the orange lines). The volcano location is indicated by a red star. (b) MISR
 695 MINX plume-height map, superposed on MISR nadir-view image. (c) MISR-retrieved plume-
 696 height profile, starting point on the east side of the MISR field-of-view. (d) Cross-track (roughly
 697 west-east) and (e) along-track (roughly south-north) MISR-retrieved wind vector components,
 698 assessed over all plume points in the digitized region. (Although this plume segment was
 699 observed by MISR to the west of the volcano, the retrieved positive zonal wind indicates
 700 eastward motion. The zonal winds varied in elevation and time on this day, and an eastward
 701 component at the location and time of the MISR observation is supported by National Center for
 702 Physical Acoustics reanalysis (NCPA, <https://g2s.ncpa.olemiss.edu/>). For all other cases in this
 703 study, the MISR-retrieved zonal wind is toward the west.) In Panel (c), Red = zero-wind height;
 704 Blue = wind-corrected height; Green = surface elevation. No wind-corrected heights were
 705 derived for the two lower layers at ~6 km and ~11 km elevation in Panel c, as the origin of
 706 digitization was selected for the more prominent layer at ~ 12 km.

707

708 **Figure 3.** MISR Research Algorithm particle property retrieval results for 15 January 2022,
 709 Terra Orbit 117440. (a) MISR context image with study region outlined; the relationship of this
 710 region to the HTHH volcano is shown in Figure 2a. (b) AOD at 558 nm in the less cloudy parts
 711 of the study region; (c) ANG assessed between 446 and 866 nm; (d) SSA at 558 nm; (e) Non-
 712 spherical AOD₅₅₈ fraction. No cost-function filter was applied in this one case, but to improve
 713 confidence in the particle property retrieval results, ANG and SSA were assessed over the AOD
 714 range $0.15 < \text{AOD}_{558} < 8.0$. Digital data for the MISR RA results are given in Table S1.

715

716 **Figure 4.** MISR MINX Plume-height retrieval results for 17 January 2022, Terra Orbit 117456,
 717 ~00:25 UTC, acquired north of Queensland, Australia, along with Lucinda AERONET ground-
 718 station retrievals. (a) MODIS context image, with MISR 60° forward view superposed (between
 719 the orange lines). The MISR retrieval region is indicated by the yellow arrow. The location of the
 720 Lucinda AERONET station on the northeast Australia coast, marked with a yellow asterisk
 721 within the MISR swath, ~450 km south of the over-ocean region where the MISR retrievals were
 722 performed, was still apparently within the plume. (For scale, the MISR swath is roughly 380 km
 723 wide.) (b) Red-band MISR-MINX plume-height histogram, showing a weak layer near the
 724 tropopause at 12-14 km and another near-surface. (c) Blue-band MISR plume-height histogram,

725 showing a layer at 27-30+ km. (d) CALIOP Level 1, version 4.51, total attenuated backscatter
726 ($\text{km}^{-1} \text{sr}^{-1}$) at 532 nm, acquired at ~05:28 UTC on 17 January, roughly 5 hours after and 800 km
727 downwind of the MISR observation at the nearest point. (The inset gives the orbit location, and
728 the red arrow highlights the likely plume layer.) (e) Lucinda AERONET AOD for 24 January
729 2022, showing the AOD peak on 17 January. (f) Lucinda AERONET AOD for 17 January 2022.
730 (g) Lucinda AERONET Ångström Exponent for 17 January 2022. (AERONET data were
731 provided by Thomas Schroeder and the AERONET team.)

732

733 **Figure 5.** MISR RA particle property retrieval results, same as Figure 3, but for 17 January
734 2022, Terra Orbit 117456. Note that the scale for panel (e) is logarithmic here to show
735 differences, given the range of in values.

736

737 **Figure 6.** MISR-retrieved Ångström exponent (ANG, inversely proportional to effective particle
738 size) for the MISR observations of the near-source HTHH plume on 15 January (red dot), and for
739 the near-tropopause (green dots) and stratosphere (blue dots and gray dot) plumes downwind.
740 Observation dates and plume heights are given in the annotations.

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