

1 **Sand, salt, and models: the legacy of Bruno Vendeville**

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17 **Keywords**

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19 Bruno Vendeville; salt tectonics; salt diapirs; analog models; sandbox models

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24 **Abstract**

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26 Bruno Vendeville (1961-2022) was the foremost practitioner of analog modeling applied to the
27 field of salt tectonics, with his work providing and promoting significant advances in our
28 understanding of this field over more than three decades. Bruno and his collaborators contributed
29 major works in: the mechanics of salt-related deformation; the key processes of extensional,
30 contractional, loading-induced, and strike-slip salt tectonics; the drivers and patterns of gravity-
31 driven, linked systems of salt-detached deformation; and the structural style and evolution of
32 specific salt basins. In this Special Issue devoted to his legacy, we offer a series of papers that
33 build on his foundational work and honor his critical role in advancing the theory and application
34 of salt tectonics.

35

36 **1. Introduction**

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38 The field of salt tectonics underwent a revolution in the late 1980s to mid 1990s. While
39 advances in understanding can be attributed in part to improved seismic reflection images, a
40 large part was provided by analog (or experimental) modeling rooted in a more realistic
41 appreciation of the mechanical behavior of layers or bodies of halite and their surrounding
42 sedimentary rocks. Although other researchers were involved, the foremost practitioner of
43 sandbox modeling during this period was Bruno Vendeville (Fig. 1).

44 This Special Issue is devoted to the memory and legacy of Bruno, who passed away in late
45 2022. We have compiled an assortment of papers dedicated to him that, like so much other
46 published research of the past 30 years, builds on the foundational work in which he was

47 instrumental. In this introduction, we highlight his key contributions to both salt tectonics and
48 analog modeling. We follow an approximate timeline of his research into salt tectonics driven by
49 extension, contraction, and differential loading, and then briefly summarize other influential
50 contributions he made during his career.

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52

53 *Figure 1. Bruno Vendeville in his laboratory at the Université de Lille in 2006 (photo by O. Ferrer).*

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57 **2. Bruno Vendeville's legacy**

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59 **2.1. The beginnings – extensional salt tectonics**

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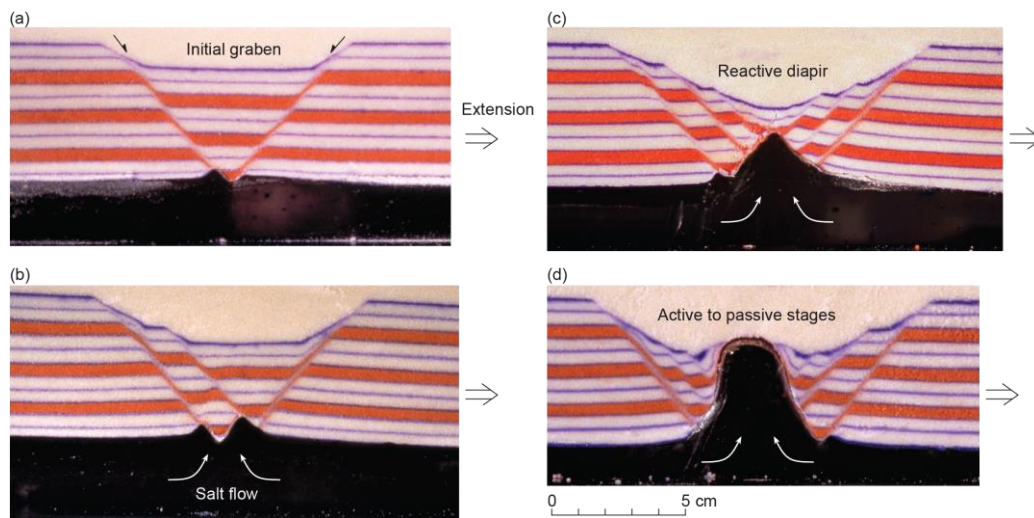
61 Bruno Vendeville undertook his doctoral studies in the mid-1980s at the University of Rennes
62 under the direction of Peter Cobbold. There he first modeled extension above a ductile
63 substratum using silicone putty as a décollement layer and layers of loose sand grains for the
64 overburden. Although some of his results were applied to crustal extension above a weak mantle
65 asthenosphere (Vendeville et al., 1987), it was the application to salt-detached extension
66 (Vendeville and Cobbold, 1987) that became the focus of his ongoing work. After receiving his
67 doctorate, he undertook post-doctoral studies at Texas A&M University before joining the late
68 Martin Jackson's Applied Geodynamics Laboratory at the Bureau of Economic Geology, The
69 University of Texas at Austin.

70 It was in collaboration with Martin Jackson that Bruno made his first major contributions.
71 Their paper on "The rise of diapirs during thin-skinned extension" (Vendeville and Jackson,
72 1992a) was a milestone. It was cited in fully half (42 of 85) of the papers contained in four
73 research volumes on salt tectonics published in the following 5 years (Cobbold, 1993; Jackson et
74 al., 1995; Travis et al., 1995; Alsop et al., 1996), was chosen as one of the top three most
75 influential papers ever published in salt tectonics by a committee of peers in 2013 (see
76 <https://100years.aapg.org/top-100-papers-salt-tectonics>), and continues to be cited widely.

77 Bruno's 1992 paper was remarkable in that it revolutionized not just one, but two, aspects of
78 salt tectonics. First, the authors argued that salt-related deformation is best viewed as a fluid
79 mechanics problem, with a pressurized fluid (ductile halite) surrounded by more competent,

80 brittle materials with significant shear strength. Bruno, in an effort to replicate this mechanical
81 contrast, therefore used sandbox models with silicone polymer as the salt analog and loose sand
82 to simulate the brittle overburden (see Weijermars et al., 1993). This was in contrast to most
83 prior models in which the overburden was treated as a higher-density viscous fluid, and
84 effectively inaugurated the “brittle era” of salt tectonics (see Jackson, 1995). Second, they
85 proposed that many diapirs are triggered by, and grow during, extension detached on the salt, and
86 they introduced the standard terminology and usage of reactive, active, and passive diapirs or
87 diapirism (Fig. 2; see Rowan and Giles, 2021). By showing that extension generates salt
88 structures, they upended the existing “halocentric” paradigm that salt movement is the driver that
89 causes normal faults to form.

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91

92 **Figure 2.** Photographs of sandbox model showing diapir initiation and growth during thin-
93 skinned extension (modified from Figs. 5 and 13 in Vendeville and Jackson, 1992a, and Fig. 3.14
94 of Hudec and Jackson, 2011): (a) – (c) reactive diapirism during extensional necking (boudinage)
95 of the brittle overburden; (d) breakthrough to the surface (active stage), where the diapir will
96 subsequently grow during the passive stage. Silicone polymer (salt) in black, sand layers in red
97 and white. Note that the arrows indicate extension, not the order of stages in the experiment.

98 Bruno's investigation of extensional salt tectonics was not confined to diapir rise during thin-
99 skinned extension. In a companion paper (Vendeville and Jackson, 1992b), he also examined
100 diapir fall during ongoing extension, emphasizing the role that weak bodies of salt exert in
101 localizing strain. Furthermore, his models elucidated the concept of partial or complete
102 decoupling during salt-involved thick-skinned extension (Jackson and Vendeville, 1994;
103 Vendeville et al., 1995), thereby setting the foundation for other researchers and explorationists
104 to model and interpret more accurately structures in such provinces as the Zechstein and
105 Pyrenean rift basins of Europe.

106

107 **2.2. Contractional salt tectonics**

108

109 Bruno soon turned his attention and talents to the problems of shortening detached on salt. He
110 first focused on the role of preexisting diapirs in localizing contractional strain due to their
111 weakness (Nilsen et al., 1995; Vendeville and Nilsen, 1995), showing how shortening can drive
112 further diapir rise (rejuvenation) even after the diapir is dead and buried (Fig. 3). He also
113 highlighted the need to look adjacent to diapirs to determine whether they had experienced
114 cryptic shortening (or, similarly, cryptic extension).

115 Again, Bruno expanded his investigation of contractional salt tectonics beyond his initial
116 contribution. Working with Elisabetta Costa, he showed that thin-skinned contraction detached
117 on salt results in different geometries and evolution than that detached on stronger, more
118 frictional shales (Costa and Vendeville, 2002). Two other influential papers addressed the role of
119 décollement dip on the geometry of accretionary wedges (Koyi and Vendeville, 2003) and the
120 impact of having two salt layers in contractional settings (Couzens-Schultz et al., 2003). In

121 addition, his models were the first to demonstrate the critical role of the preexisting three-
122 dimensional diapir and minibasin distribution in controlling the subsequent development of
123 polygonal contractional structures (Rowan and Vendeville, 2006). This proved foundational in
124 the reinterpretation of numerous orogenic belts with unusual structural patterns due to the
125 presence of salt and an early history of diapirism.

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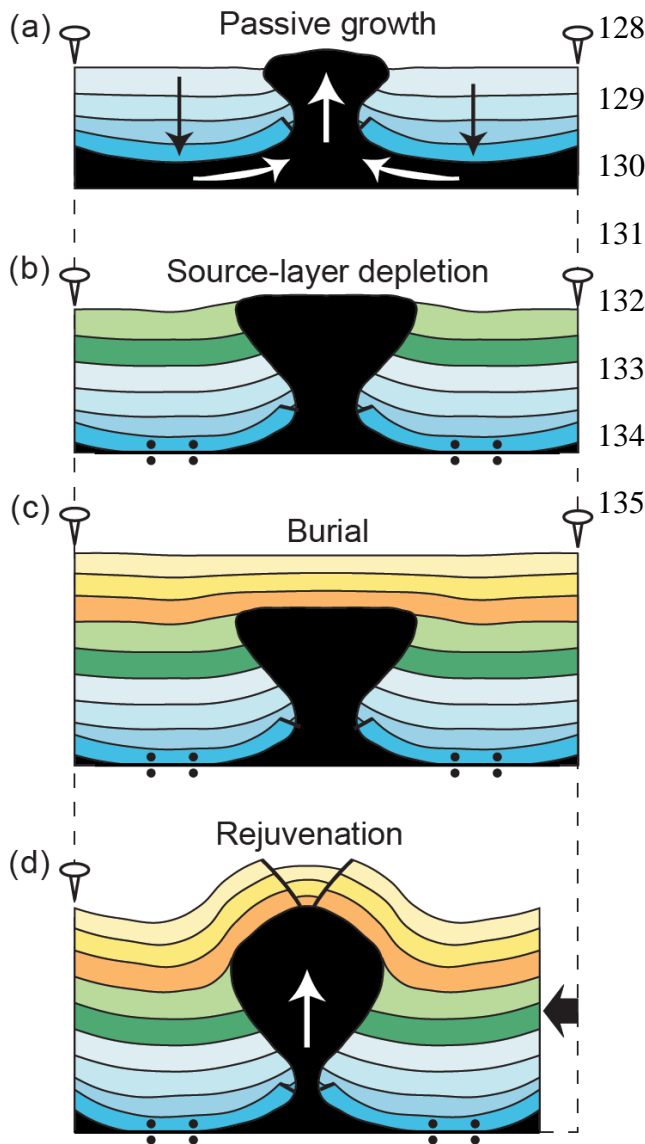


Figure 3. Sketch of model results showing burial of a passive diapir due to source-layer depletion and then contractional rejuvenation (modified from Fig. 12 in Vendeville and Nilsen, 1995, and Fig. 2.87 in Hudec and Jackson, 2011). Silicone polymer (salt) in black, pairs of dots denote salt welds.

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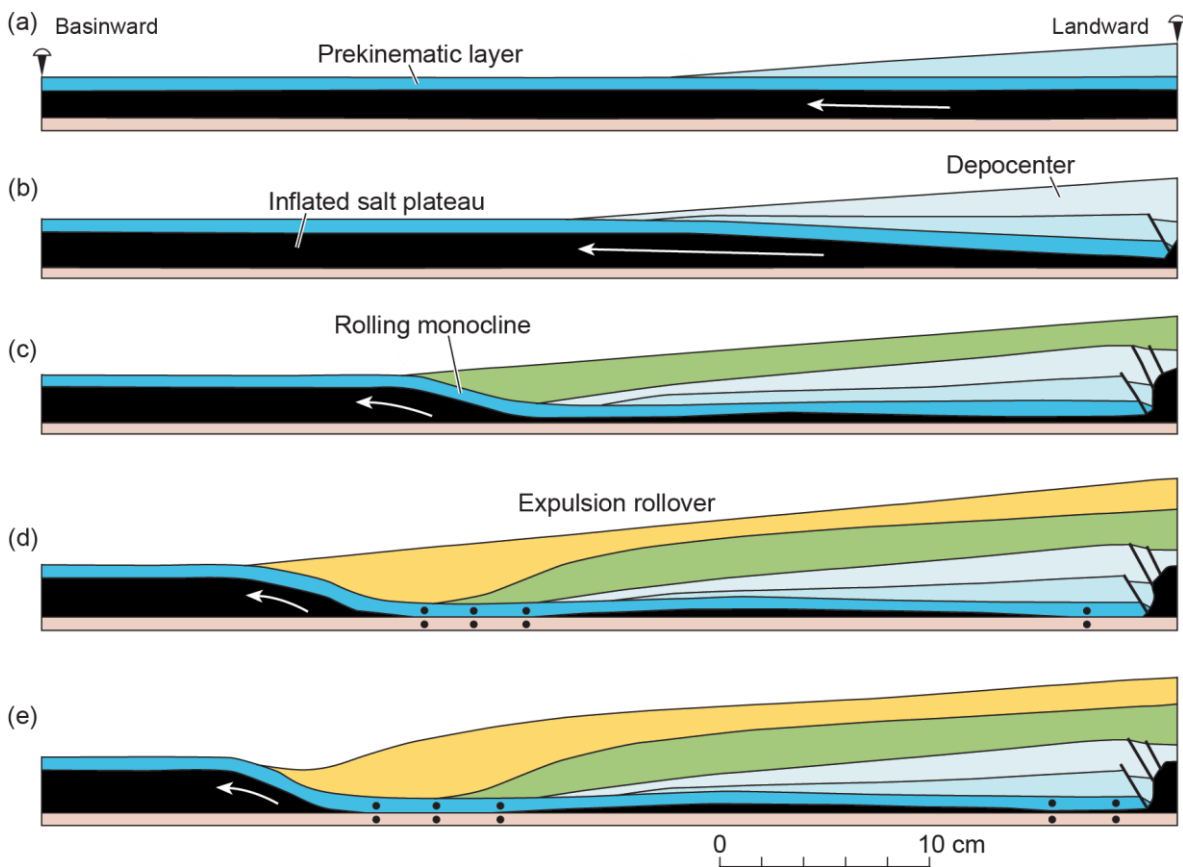
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138 **2.3. Loading-driven salt tectonics**

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140 Bruno also conducted investigations of salt tectonics in the absence of either extension or
141 contraction. In an influential paper (Ge et al., 1997), he and colleagues modeled progradational
142 loading above salt, showing that the differential load was sufficient to drive salt basinward (Fig.
143 4) regardless of the density contrast between salt and its overburden, producing salt pillows,
144 walls, and sheets. This model has subsequently been applied to many salt basins such as the
145 Southern Permian Basin, the northern Gulf of Mexico, and the Paradox Basin.

146



147

148 **Figure 4.** Sketch of model results showing progressive expulsion of salt to produce basinward-shifting
149 depocenters known as expulsion-rollover structures (modified from Fig. 6 in Ge et al., 1997, and Fig.
150 1.32 in Hudec and Jackson, 2011). Silicone polymer (salt) in black, pairs of dots denote salt welds.

151 Bruno pursued these ideas in two further papers on the mechanics, kinematics, and three-
152 dimensional geometries of linked systems of proximal extension, medial translation, and distal
153 contraction generated by gravity spreading of depositional wedges above salt (Vendeville, 2005;
154 Gaullier and Vendeville, 2005). This tied in to several papers on the triggers, drivers, and styles
155 of gravity-driven deformation on rifted margins (Rowan et al., 2004, 2012). Bruno provided the
156 mechanical reasoning and supporting analog models in arguing that gravity spreading played an
157 important, and sometimes dominant, role in driving the deformation on several margins, in
158 contrast to models claiming that only gravity gliding drives gravitational failure. Once again,
159 Bruno's work proved foundational for all subsequent studies of gravity-driven deformation on
160 rifted margins such as the Gulf of Mexico and South Atlantic salt basins.

161

162 **2.4. Other significant contributions**

163

164 Bruno continued publishing influential papers after his move to the Université de Lille in
165 2004, where he remained for the rest of his career. He conducted and published further research
166 on extensional (e.g., Ferrer et al., 2014), contractional (e.g., Santolaria et al., 2015; Borderie et
167 al., 2018), and loading-induced salt tectonics, but also on other aspects such as strike-slip salt
168 tectonics (e.g., Boussarsar et al., 2022; Vendeville et al., 2024). His laboratory at Lille also
169 contributed a series of important papers on various aspects of deformation above the Messinian
170 salt of the Mediterranean (e.g., Gaullier et al., 2000; Loncke et al., 2006). In addition, Bruno
171 tackled non-salty topics such as shale tectonics (e.g., Mourgues et al., 2009) and the role of
172 overpressured fluids in deformation (e.g., Lacoste et al., 2011). In all cases, his contributions
173 were informed by innovative analog models.

174 **3. Papers in this volume**

175

176 Bruno's legacy lies not just in his papers and the many new ideas that were tested and refined
177 by experimental modeling. He also stimulated and inspired others who followed in his footsteps
178 in their own modeling labs, including, but not limited to: Tim Dooley, who took over Bruno's
179 position in the Applied Geodynamics Laboratory in Austin; Jean-Paul Callot, first at the IFP in
180 Paris and then the Université de Pau; Oriol Ferrer and colleagues at the Geomodels lab in the
181 Universitat de Barcelona; and Jürgen Adam at Dalhousie University and then Royal Holloway
182 University of London.

183 In the following section, we briefly summarize the fifteen papers published in this Special
184 Issue of the Journal of Structural Geology. These papers originate from some of the labs
185 mentioned above, other researchers, and Bruno's own lab at Lille. Notably, Bruno is an author of
186 five of the contributions. In organizing the papers, we somewhat loosely follow the general
187 progression of Bruno's own work over the decades. We thus start with extensional salt tectonics
188 and then move on to contractional, strike-slip, gravitational, and loading-induced salt tectonics,
189 finishing with a paper on intrasalt deformation.

190 The special issue begins with Schöpfer and Lehner (2024), who take a new look at Bruno's
191 most famous modeling results – the rise and fall of diapirs during thin-skinned extension – but
192 from a numerical modeling perspective. They reproduce many of the salt structures originally
193 generated in the sandbox models of Vendeville and Jackson (1992a, 1992b), including reactive,
194 active, passive, and collapsed diapirs, turtle and mock-turtle structures, and salt horns.
195 Interestingly, they demonstrate that diapir fall during late-stage extension is dependent on the
196 shape of the diapir (triangular, plug-shaped, or hourglass-shaped). Moreover, both turtle and

197 mock-turtle structures in their models of thin-skinned extension generally have salt cores rather
198 than underlying welds.

199 The paper by Evans et al. (2024) uses analog models to investigate the interplay of base-salt
200 relief, salt thickness, and intrasalt brittle layers during gravity-driven deformation in the
201 extensional domain. Key findings are that: (i) a first-order control is the thickness of salt relative
202 to the magnitude of base-salt relief; (ii) thin salt promotes the development of ramp-syncline
203 basins, lateral over vertical movement, and stronger spatial coupling between supra- and subsalt
204 deformation; (iii) thick salt leads to reactive diapirs, extensional rollover and turtle structures,
205 more vertical motion, and weaker coupling; and (iv) intrasalt heterogeneity increases the degree
206 of coupling, inhibits vertical motion, and promotes distributed normal faulting.

207 The Special Issue includes more papers on contractional salt tectonics than on other topics.
208 We start with the paper by Santolaria et al. (2024), who investigate the origins of oblique
209 structures in fold-and-thrust belts. They first synthesize previously published analog models
210 using linear backstops but laterally varying cover and/or décollement thicknesses, discussing
211 different material properties, model setups, tested parameters, and consequent results. They then
212 present their own analog models in which they examine gradually varying lateral thickness
213 changes in both polymer (salt) and cover, specifically scenarios where the two thin in the same
214 or opposite directions. They revisit two natural case studies (the South Pyrenean Central Salient
215 and the Keping Shan fold-and-thrust belt of northwestern China) in light of the various model
216 results, concluding that analog models are invaluable in guiding interpretations but cannot
217 provide unique solutions.

218 The paper by Muñoz et al. (2024) also uses analog models to investigate the origin of oblique
219 structures and explain the South-Central Pyrenean Thrust Salient. Their models employ

220 trapezoid-shaped detachments of silicone polymer surrounded by frictional material and overlain
221 by foreland-thinning wedges of sediment. The most informative models have two similarly
222 oriented but non-parallel lateral edges of silicone, such that one side has a laterally widening
223 frictional area between the backstop and a narrowing area of silicone, and the other has a simple
224 laterally narrowing silicone detachment. They show that the salient is largest at early stages of
225 deformation but becomes less prominent over time, that the viscous and frictional thrust wedges
226 have different kinematics and styles and are separated by transition zones, and that oblique
227 structures in these zones rarely have the same trends as the original edges of the silicone.
228 Moreover, along-strike variations in model wedge taper and consequent topography nicely
229 explain the asymmetric relief of the Pyrenean salient, the development of the Ainsa Basin, and
230 the attendant routing of synorogenic sediment.

231 Feng et al. (2024) employ analog models to investigate the influence of mechanical
232 stratigraphy and décollement thickness on the geometry and kinematics of salt-detached fold-
233 and-thrust belts, with specific application to the Kuqa Basin of China. Thinner salt leads to the
234 development of thrust faults with minimal early folding, whereas thicker salt favors the
235 formation of break-thrust folding and eventually source-fed thrust faults. Including layers of
236 mica flakes creates a layer anisotropy and an initially stronger overburden, but ultimately weaker
237 and thus longer-lived, higher-displacement thrust faults. Both thicker salt and mica flakes result
238 in more bulk movement of salt toward the foreland and the consequent growth of inflated salt
239 massifs.

240 Peng et al. (2024) also use analog models to guide interpretation of structures in China,
241 specifically the interaction between inherited basement strike-slip structures and thrust-wedge
242 propagation in the northern Tianshan foreland basin. Although the primary sedimentary

243 décollement is overpressured shale, the results are also applicable to salt detachments because
244 the authors use silicone polymer as the viscous layer. They show that folding of the detachment
245 by reactivation of the basement structure prior to contraction, especially if accompanied by
246 erosional thinning of the overburden, leads to strain localization over the deeper structure and
247 out-of-sequence thrusting.

248 Célini et al. (2024) refer to published analog models to help explain and understand structural
249 styles in the external Western (French) Alps by invoking along-strike variations in the presence
250 and thickness of Triassic salt. Where there was no salt (Northern Subalpine Chains), the
251 structural style matches models with frictional detachments. Where the salt was relatively thin
252 and in a foreland position (Jura Mountains), the style changes from a frictional wedge in the
253 hinterland to a viscous wedge in the foreland, as seen in models. Where the salt was thick (SW
254 Alps), the structural styles resemble those of analog models with preorogenic diapirs and
255 minibasins, with features such as squeezed diapirs, overturned megaflaps, salt sheets, and
256 secondary minibasins. Finally, the spatial distribution of salt controlled the development of
257 salients and reentrants along the Western Alps wedge.

258 In the last paper on contractional salt tectonics, Krzywiec et al. (2024) focus on thick-skinned
259 inversion, using both analog and numerical models to examine the genesis of seismically imaged
260 salt pillows in central Poland. The models are mutually consistent and show that pillows form
261 above inverted subsalt half-graben containing originally thicker salt, with the most important
262 variables influencing the final geometry being the size and geometry of the half-graben, the
263 viscosity of the salt, and the amount of any erosional thinning of the overburden during pillow
264 formation. The results have significant implications for depositional thickness and facies
265 reconstructions of the Zechstein and other synrift salt basins.

266 We then move to a historical paper on strike-slip salt tectonics by Vendeville et al. (2024).
267 Bruno carried out these models in the late 1990s but never had them published. Although many
268 more recent models on strike-slip deformation have come out in the intervening years, this
269 methodology paper illustrates Bruno's innovative approach to analog modeling. Instead of
270 generating strike-slip deformation in model sand by the movement of rigid underlying plates,
271 Bruno builds two weak zones within the sand above a layer of silicone polymer before applying
272 end-wall shortening. These are intended to represent preexisting salt structures in thin-skinned
273 deformation or established crustal basins during basement-involved shortening. Localized
274 contractional strain in the weak zones is linked by different styles of strike-slip structures
275 depending on the model setup.

276 An example of this approach is provided by Zidi et al. (2024). They examine an area of the
277 Tunisian Atlas Mountains where there are outcrops of Triassic salt along a fault system with two
278 segments at about 45° to each other. They postulate the existence of an original buried salt wall
279 with this kinked geometry and use analog models to test its reactivation during contraction.
280 Whereas the segment oriented perpendicular to the shortening direction is squeezed orthogonally
281 until a vertical weld forms, the oblique segment is reactivated with strike-slip movement, and
282 areas without any preexisting diapir develop simple folds and thrust faults. Local extrusions of
283 salt occur along both segments of the structure, matching the observations from the natural
284 prototype.

285 The Special Issue includes three papers on large-scale linked systems of gravity-driven
286 deformation. First, Anagnostoudi et al. (2024) employ complex analog models in an attempt to
287 reproduce the salt-related geometries of the Levant Basin in the eastern Mediterranean. They
288 model four scenarios – gravity gliding/spreading from the linear to slightly curved Levant

289 margin, gravity spreading from the highly arcuate Nile deep sea fan, and two different
290 combinations of both drivers –including in each the distal buttressing effects of the Eratosthenes
291 Seamount. Although no single model fully reproduces the three-dimensional patterns of the
292 Levant Basin, they conclude that gravity spreading of the Nile cone was the dominant driver of
293 deformation, with relatively minor contributions of gliding/spreading from the Levant margin.

294 The paper by Travan et al. (2024) is the first of two that examine salt tectonics driven by a
295 combination of tectonic shortening and gravitational failure. They use a sandbox model to
296 evaluate the genesis and evolution of salt structures in the offshore of the central segment of the
297 Algerian Mediterranean margin. Local uplift of the Messinian salt and its overburden was
298 generated by underlying contractional deformation due to convergence between the African and
299 European plates. The model suggests that this in turn led to along-strike gravity gliding down the
300 flank of the plateau, with bulk movement of salt downdip to the west and the development of
301 both extensional structures and a polygonal array of diapirs and minibasins due to convergent
302 shortening.

303 Dooley and Hudec (2024) focus on the Sureste Basin in the southern Gulf of Mexico,
304 progressively building complexity into analog models to address five aspects of the observed
305 geology. They demonstrate that: *(i)* basinward narrowing of a trapezoidal salt basin results in
306 constrictional strain and transpressional deformation along both lateral margins; *(ii)* a
307 combination of tectonic and gravity-driven deformation is required to produce contractional
308 structures throughout the entire basin; *(iii)* the presence of preexisting diapirs results in
309 curvilinear and kinked fold and thrust-fault structures linking the diapirs; *(iv)* non-parallel
310 driving forces generates strong curvature of structural trends and major transpression along the

311 western boundary of the basin; and (v) an eastward-thinning salt layer explains the distribution of
312 salt sheets and canopies in the Sureste Basin.

313 Ferrer et al. (2024) bring us back to more local deformation in the first of two final papers,
314 investigating the development of halokinetic megaflaps adjacent to passive diapirs. They use
315 analog models to investigate differential loading with either vertically-stacked depocenters or
316 laterally-shifting expulsion rollovers. In both cases, there is earlier subsidence of one minibasin
317 relative to another; touchdown and consequent welding of the salt beneath the first then leads to
318 increased subsidence, stratal rotation, and megaflap formation in the later minibasin. Moreover, a
319 more cohesive mixture of sand and clay in the prekinematic layer results in minimal bed
320 lengthening and thinning, compatible with observations of exposed megaflaps.

321 In the final paper, Cofrade et al. (2024) examine the distribution, geometry, and evolution of
322 intrasalt carbonate stringers in an allochthonous salt tongue at Les Avellanes diapir in the south-
323 central Pyrenees. They employ an analog model to demonstrate that: (i) stringers are rotated
324 from near-vertical in the feeder to subhorizontal in the salt sheet; (ii) intra-sheet deformation
325 generates tight to isoclinal, inclined to recumbent folds, vergent toward the sheet toe, that
326 become disrupted and stacked with ongoing flow; and (iii) the stringers become imbricated
327 toward the toe of the sheet due to decelerating flow and consequent compressional stress.

328

329 **4. Concluding statement**

330

331 We hope and expect that this assortment of papers offers a fitting tribute to the memory of
332 Bruno Vendeville. We stress that it is not only his scientific contributions that we honor, but also
333 Bruno the friend, colleague, and mentor. He truly was salt of the earth.

334

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336

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347

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