

BUREAU OF ECONOMIC GEOLOGY STATE OF TEXAS ADVANCED RESOURCE RECOVERY and PETROLEUM TECHNOLOGY TRANSFER COUNCIL

Present

Sequence Stratigraphy, Depositional Systems, and Production Trends in the Atoka Series and Mid-Pennsylvanian Cleveland and Marmaton Formations, Western Anadarko Basin

Lipscomb

DAVID L. CARR, TUCKER F. HENTZ, WILLIAM A. AMBROSE, ERIC C. POTTER, AND SIGRID J. CLIFT

Schil

TUESDAY, NOVEMBER 10, 2009 ELLISON MILES INSTITUTE Brookhaven College, Farmers Branch, Texas 8:30 a.m. – 4:00 p.m.

BUREAU OF ECONOMIC GEOLOGY Scott W. Tinker, Director Jackson School of Geosciences The University of Texas at Austin

No. SW0019

QAe7466



BUREAU OF ECONOMIC GEOLOGY STATE OF TEXAS ADVANCED RESOURCE RECOVERY and PETROLEUM TECHNOLOGY TRANSFER COUNCIL

Present

Sequence Stratigraphy, Depositional Systems, and Production Trends in the Atoka Series and Mid-Pennsylvanian Cleveland and Marmaton Formations, Western Anadarko Basin

David L. Carr, Tucker F. Hentz, William A. Ambrose, Eric C. Potter, and Sigrid J. Clift

TUESDAY, NOVEMBER 10, 2009 Ellison Miles Institute Brookhaven College, Farmers Branch, Texas 8:30 a.m. – 4:00 p.m.

PTTC gratefully acknowledges support of industry, academia, and the DOE's National Energy Technology Laboratory. This material is based upon work supported by the Department of Energy under Award No. DE-FE0001175.

> **BUREAU OF ECONOMIC GEOLOGY** Scott W. TINKER, DIRECTOR JACKSON SCHOOL OF GEOSCIENCES THE UNIVERSITY OF TEXAS AT AUSTIN

> > No. SW0019

Workshop Agenda

Morning Session – 8:30 a.m.-11:20

- I. Introduction and Welcome *Eric Potter*, Bureau of Economic Geology
- II. Atoka Series, Play Overview David Carr, Bureau of Economic Geology

Break – 10:15-10:30

III. Sequence stratigraphy and depositional fabric of the Marmaton and Cleveland Formations *Tucker Hentz*, Bureau of Economic Geology

Lunch

We will break for lunch at 11:20 for those who are attending the DGS luncheon meeting. Workshop will convene again at 1:00.

For those who are not attending the DGS luncheon meeting, lunch will be provided in the workshop classroom.

Afternoon Session – 1:00 p.m.-4:00 p.m.

- IV. Tidally dominated depositional systems and sedimentary processes, with application to the Marmaton and Cleveland Formations *Bill Ambrose*, Bureau of Economic Geology
- V. Core workshop Bill Ambrose, David Carr, Tucker Hentz

Instructor Profiles

William A. Ambrose is a geologist specializing in sedimentology and reservoir characterization. He received a M.A. degree in geological sciences in 1983 from the University of Texas at Austin. His contact information is--email: william.ambrose@beg.utexas.edu , telephone: 512-471-0258.

David L. Carr is a geologist whose interests lie in clastic sedimentology and stratigraphy and their application to exploration, production, and gas-storage activities. He earned his M.A. degree geological sciences in 1983 from the University of Texas at Austin. His contact information is: email – <u>david.carr@beg.utexas.edu</u>, telephone – (512) 471-1806, address – Bureau of Economic Geology, The University of Texas at Austin, University Station, Box X, Austin, TX 78713-8924.

Tucker F. Hentz is a geologist with the Bureau of Economic Geology specializing in sequence stratigraphy and basin analysis. He received his M.S. degree in geology in 1982 from The University of Kansas. His contact information is: email – <u>tucker.hentz@beg.utexas.edu</u>, telephone – (512) 471-7281.

Eric Potter is Associate Director of the Bureau of Economic Geology and is responsible for managing the Bureau's energy-related research. Eric received a B.A. degree in Geology at Dartmouth College in 1972 and a M.S. degree in Geology at Oregon State University in 1975.

Sequence Stratigraphy, Depositional Systems, and Production Trends in the Atoka Series and Mid-Pennsylvanian Cleveland and Marmaton Formations, Western Anadarko Basin

Workshop Presented by the Bureau of Economic Geology State of Texas Advanced Resource Recovery (STARR) Program and PTTC Texas and SE New Mexico Region

> November 10, 2009 Ellison Miles Geotechnology Institute, Dallas

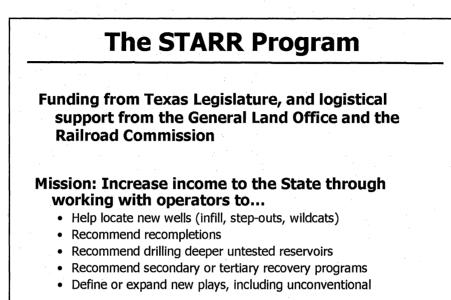
Bureau of Economic Geology

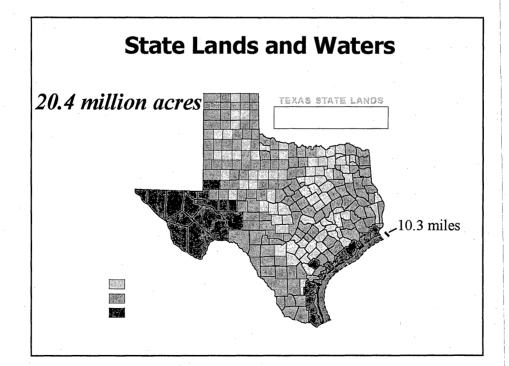


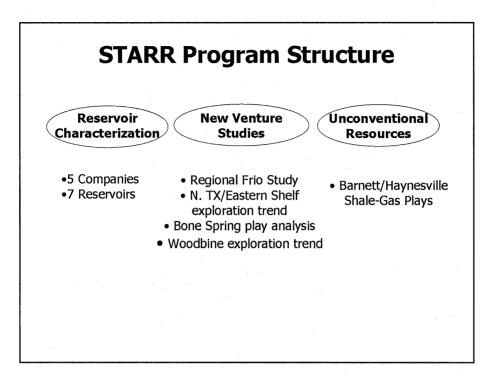


Introduction

Eric Potter BEG Program Director for Fossil Energy







STARR Partnerships with Companies

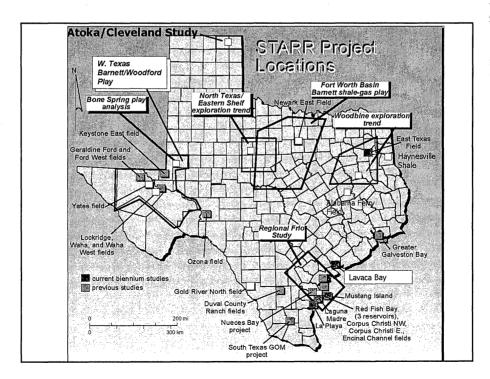
- No cost to operators
- Publication of some results

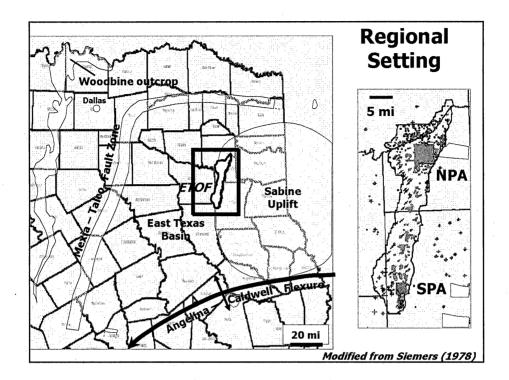
BEG's STARR partner selection criteria:

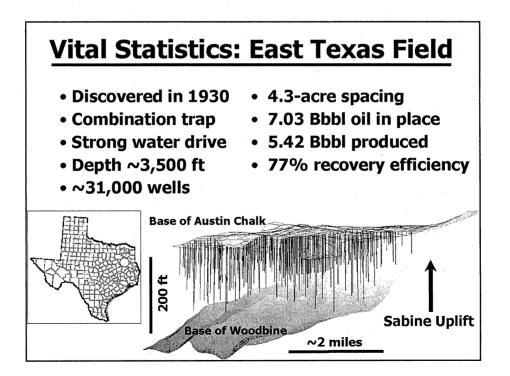
- Potential economic value to Texas
- · Operator's willingness to share data
- Operator's financial ability to drill, recomplete, and/or initiate an enhanced recovery program
- Willingness of operator to write letter describing outcome of STARR recommendations

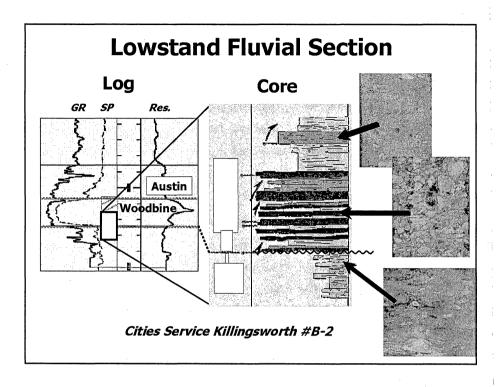
STARR Team Members

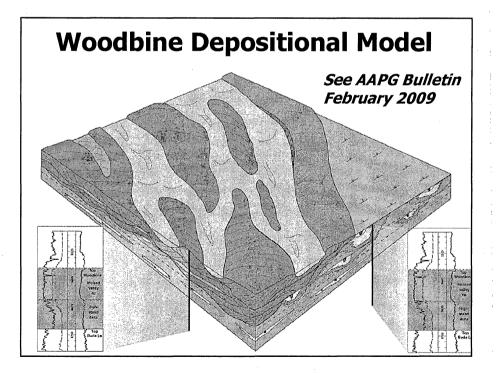
Ursula Hammes, co-PI Bill Ambrose, co-PI Frank Brown (sequence stratigrapher) Florence Bonnaffe (geologist) Cari Breton (GIS) David Carr (geologist) Ray Eastwood (petrophysicist) Julia Gale (structural geologist) Tucker Hentz (geologist) Brandon Johnson (geophysicist) Lorena Moscardelli (seismic interpreter, geologist) Chris Ogiesoba (geophysicist) Fred Wang (petroleum engineer) Hongliu Zeng (geophysicist) Scott Hamlin (geologist)

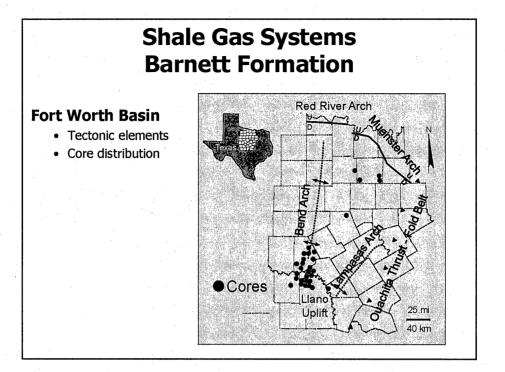


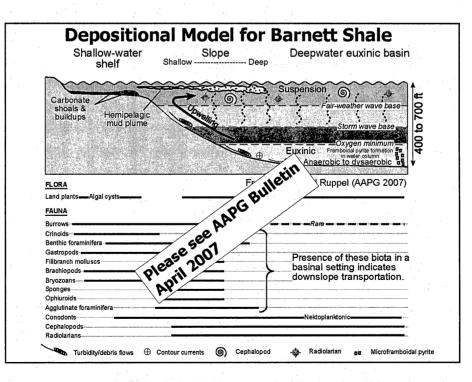












STARR Financial Model

- STARR Project must be revenue neutral to be re-funded by the Legislature each biennium
- Presently funded at \$1.5 million annually
- STARR return on State's investment was >6:1 in last biennium, despite stringent credit criteria
- Partnerships with companies are critical to program's success
- Remember, it's free!

Agenda for today's workshop

Morning Session (8:30 a.m.-Noon)

- Atoka Series
 -Play Overview (David L. Carr)
- Cleveland Formation and Marmaton Group -Sequence Stratigraphy (Tucker F. Hentz)

Lunch (Noon-1:00 p.m.)

-Optional Luncheon (Jerry Lucia, speaker)

Afternoon Session (1:00 p.m.-4:00 p.m.)

Cleveland Formation and Marmaton Group
 -Tidal Depositional Systems (William A. Ambrose)
 -Core Workshop (Carr, Hentz, Clift, Potter, Ambrose)

Atoka/Cleveland Team



Dave Carr



Tucker Hentz







9

Atoka and Cleveland/Marmaton Plays

• Atoka Series Variable rock types: limestone and shale High-TOC shales (avg. 4-6%; some >10%)

• Cleveland Formation and Marmaton Group

Low-permeability sandstones High degree of interbedded sandstones and shales Narrow sandstone-body geometries Partitioning of sandstones in sequences

Databases

Atoka Series

>500 Wells Shell #1 Molesworth core – 84 ft Production data from 525 vertical and 20 horizontal wells

Cleveland Formation and Marmaton Group

>1,100 Wells Cores from 7 wells Production data from >900 wells

Production Characteristics

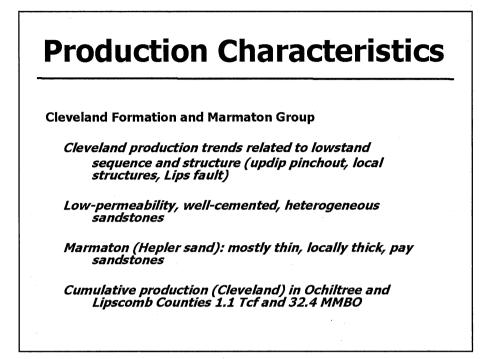
Atoka Series

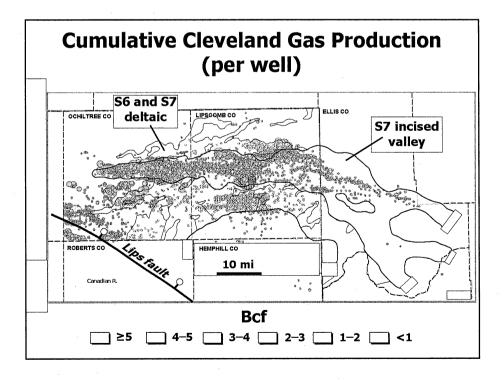
Main production: Thirteen Finger Ls-variable quality

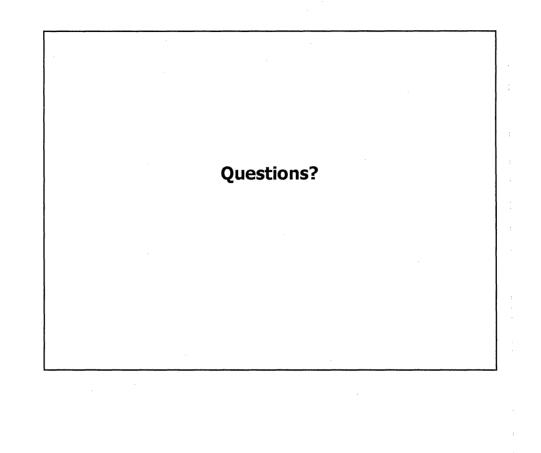
Cumulative production in Ochiltree and Lipscomb Counties 185 Bcf and 6.7 MMBO

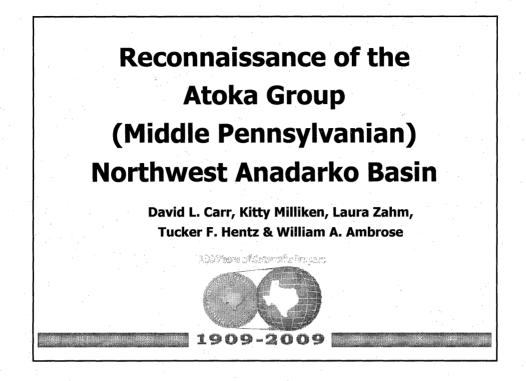
Potential for unconventional shale-gas production

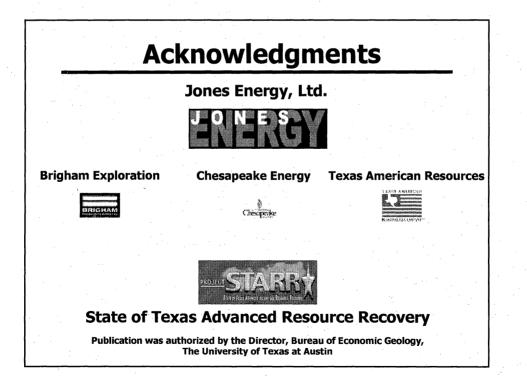
Some structural control on production trends

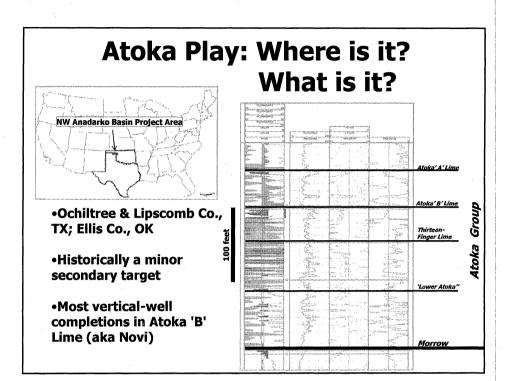








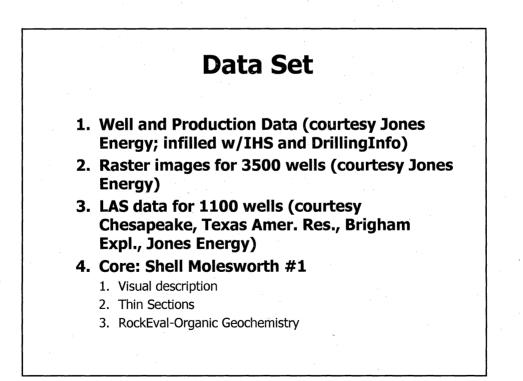


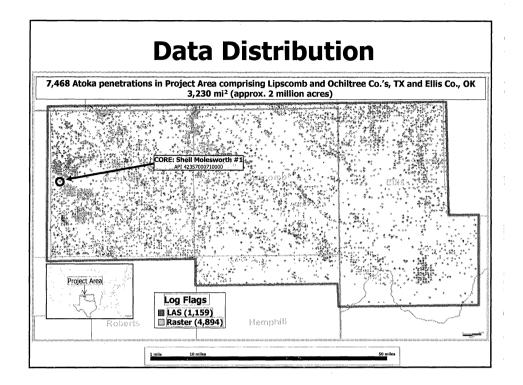


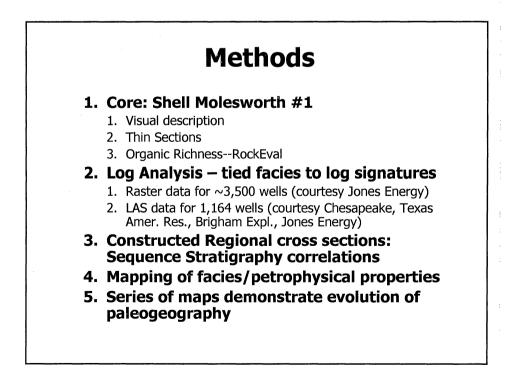
<section-header><section-header><text><list-item><list-item><list-item><list-item><list-item><list-item><list-item><list-item><list-item><list-item><list-item>

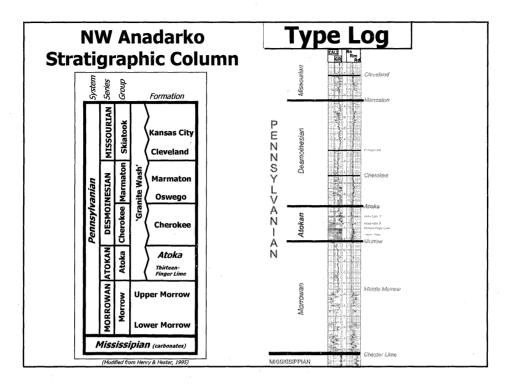
Outline

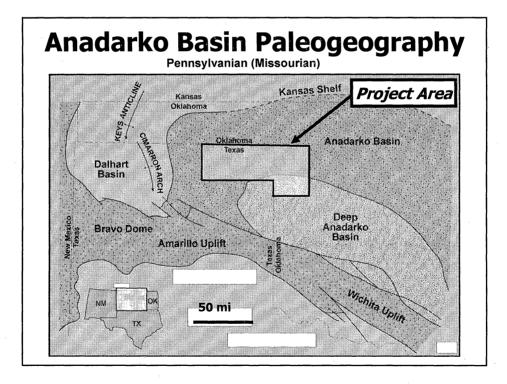
- 1. Intro: Data, Methods & Geologic Setting
- 2. Production History & Current Activity
- 3. Core Calibration: Shell Molesworth #1
 - 1. Facies
 - 2. Inferred Depositional Environments
 - 3. Organic Richness
 - 4. Petrophysical Properties
- 4. Sequence Stratigraphy
- 5. Facies Mapping/Paleogeography

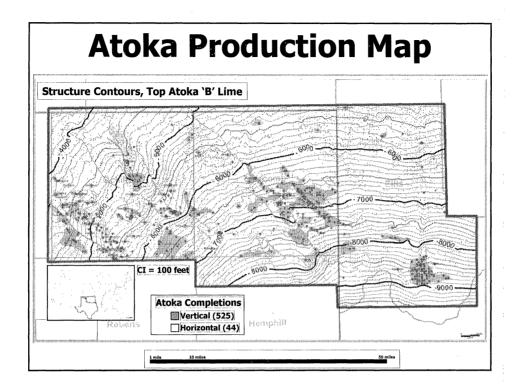


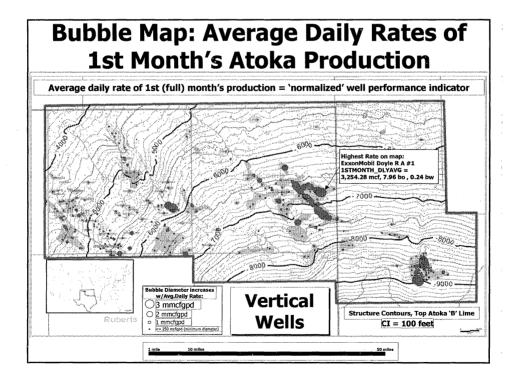












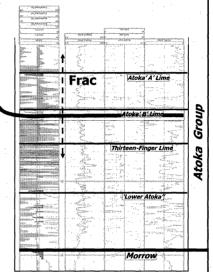
Emerging Atoka Horizontal Lime/Shale Play

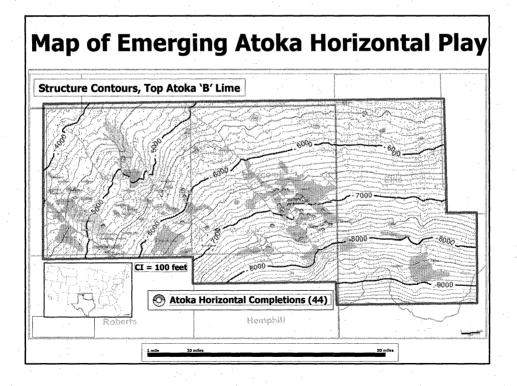
•Historically a minor secondary target

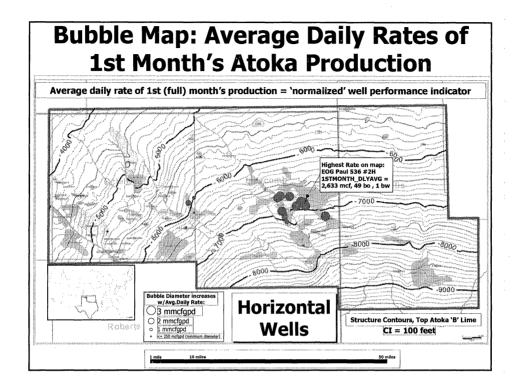
•Most vertical-well completions in Atoka 'B' Lime (aka Novi)

•Emerging horizontal/ fracture stimulation play targeting the Atoka 'B' Lime

How much of this increase is the result of hydraulic fractures reaching into the adjacent organic-rich shales?







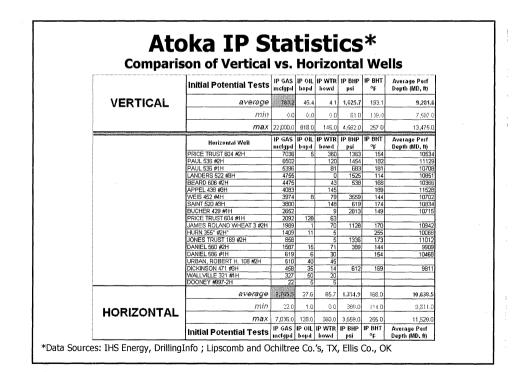
~~~~~	Cumulative Production Average Daily Rates												1
	GAS (mcf)	OIL (bbl)	WATER (bbl)		GAS max rate mcfgpd		OiL avg rate bond	OIL max rate bopd	OiL min rate bopd	WTR avg rate bwpd	WTR max rate bwpd	WTR min rate bwpd	Days on Prod.
31997	219,463,496			-		-	*	*		-	*	*	1,601,448
everage	417,231	20,618		180.4	575.8	45.9	13.7	1,095.8	122.4	2.3	87.4	11.1	3,048
(MAX	17.316.336	1,674,521	102,078	4,215,2	\$88.9	48,8	1,000.5	1.419.2	1,019.6	100.0	140.3	59,1	17.017

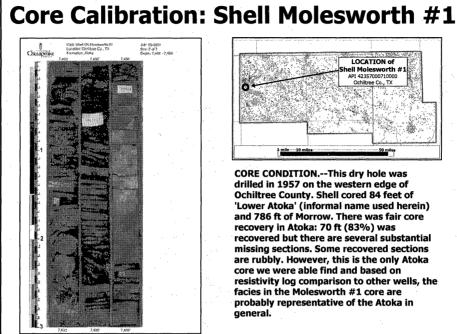
	Horizontal Wells (n = 21)														
			Consulative Production			Augustia Daily, Paton								1	
Viell	10	Compt. Date	GAG	08.	WATER Dbi				08. 379 7857	08, 1983 1989	OR. ssice reats	WTR avg rate boypd	WTR max calk bwpd	WTR 989 (als bapd	Days o Prod.
Dath KAAmma	1 30 8080	8.022.06	4285 254	125 125	7 236	2.832	4.900	\$ 1834			<u></u>	30	30	28	15
LANCERS 522 #3H			315.4611	24	(		4.549	885	8.85	ž		1	· · · · · ·		1
Appel, 438 #3H			168,457	285	*	2.668	2,670	2.240	4.68	ŝ	4		•	•	
5a847 520 #3H	10.000	28-44-08	194,474		•	\$,698	2,826	5.002		•	~				
Ранц 336 фін				201	7,162	1,494	4.20%	720	1,84	7	õ	21	61	10	22
Price trust 604 Kom	\$1.666	24-540-01	638,848	287	21.228	5,493	4.004	5.62	0.73	3	8	\$0.	\$83	2.2	28
NEIS 452 #4H											3			11	
Brard 605 \$2H				\$\$.				3.46			<u> </u>	14			48
											<u>0</u>				
									9.53			{?! <u>*</u>			15 18
				······`ya				ş				þ			
											·····;%:		· · · · · ·	·····	
					ŀ							§			او
					l sûar			<u>.</u>		in the second			·····		
			122 8 22 1		2 806	158		28	1.50			for a start of the	8 8	1	
JONES TRUST 169 40H										······	نیک B				 53
WALLVILLE 321 #1H			24.278		1.284	48	127			.22	Q.	2	3	0	
GOONEY #967-2H	0.705	24-Oct-05	4,781 [	830	224	5	\$ }	§ §		21	Q.	1	2	9	88
COURSON RANCH 136 #13H	3,478	3.1ter-08	188			1	E		3.23					uuuuiiuuu	
		aunu	5,009,646	32,099	75,484	×14	ine .	.89	ne	712	209	.a	60	15.6	8.0
														e	3
			230.126	1,687	3.594	\$43.9	1,842.8	517.2 2.239.8	4.2	30.7	2.1	19.3	28.9	5.6	
	CAUL 1020 MICH CAUCERT 5/2 4944 SAPEL 318 SCH SAPEL 300 RISH PRILICE TRUST 300 RISH PRILICE RISH 300 RISH PRILICE R	Sale         Sale         Sale           CAUGERS (S2 HM1         10.050           CAUGERS (S2 HM1         10.050           Sale (S3 MAR)         11.050           Sale (S4 M	Verti         U         Date           PAUL 500 WSN         12.559         Skept 28.           Aud CRED 52.2541         17.069         Skept 28.           Aud CRED 52.2541         10.000         Skept 28.           Skept 28.3641         10.000         Skept 28.           Skept 28.3641         11.000         Skept 28.           Skept 28.         11.061         Auton 27.           Skept 28.         11.061         Skept 28.           Skept 28.         Skept 28.         Skept	Viell         TO         Compt. Date         G&G           PAUL 5:01 WCH         12:559         6-087-081         402.014           AUGRER 5:22:8514         10:000         15.46-01         564-061           AUGRER 5:22:8514         10:000         15.46-01         164.021           AUGL 5:000 HL         11:078         6-564.07         64.024           AUGL 5:00 HCM         11:058         24.000         553.462           VIES 5:22:82:82         11:058         24.000         37.000           DEAHER 3:00 HL         11:058         24.000         37.000           DEAHER 4:00 HL         11:058         24.000         37.000           DEAHER 4:00 HL         11:058         24.000         37.000           DEAHER 4:00 HL         11:058         11:062         38.000           DEAHER 4:00 HL         11:078         14.000         38.000           DEAHER 4:00 HL         10:078         11:082         38.000           DEAHER 4:00 HL         10:0	Weft         TO         CompA Date         G&0         Oil.           PAUL 550 KRN1         12,559         5-692,05         402,816         568         568           AUGERD 522 RM1         10,000         15-468         768         568,452         225           AUGERD 522 RM1         10,000         15-468         768         568,452         225           AUGERD 522 RM1         10,000         15-468         768         568,452         225           AUGERD 522 RM1         10,000         12-468         568,452         225           AUGERD 502 RM1         11,000         444,692         268         268           VIES 52,000 M         11,058         44469,07         268,652         215,660           VIES 52,000 M         11,058         44469,07         268,652         215,660           VIES 52,000 M         11,058         14469,07         268,652         215,660           VIES 52,000 M         11,058         14469,07         216,862         216,860           VIES 52,000 M         11,058         14469,07         216,862         216,860           VIES 52,000 M         10,978         11,982         126,862         216,862           VIES 54,000 M         10,978 <td>Vietl         TO         Date         GAR         DBI           FAUL K50 wgAr         12,549         6-bgr-26         442,754         136         7.42           AUGERDS K22 #341         10,000         5-fam-26         344,457         7.43         7.44           AUGERDS K22 #341         10,000         5-fam-26         154,467         2.25         -           AUGERDS K22 #341         10,000         2-fam-26         164,474         -         -           AUGERDS K22 #341         10,000         2-fam-26         164,474         -         -           AUGERDS K22 #341         10,000         2-fam-26         164,427         -         -           AUGERDS K22 #344         10,000         2-fam-36         164,427         -         -           VESI 50,242         2-fam         164,612         2-fam&lt;27</td> 561,480         2.14         -           VESI 50,242         2-fam         164,612         -         -         2.14         -         -           VESI 50,410         2-fam         569,677         174,122         -         -         -         -         -           VESI 50,410         VESI 50,540         11,1750         554,967         114,467	Vietl         TO         Date         GAR         DBI           FAUL K50 wgAr         12,549         6-bgr-26         442,754         136         7.42           AUGERDS K22 #341         10,000         5-fam-26         344,457         7.43         7.44           AUGERDS K22 #341         10,000         5-fam-26         154,467         2.25         -           AUGERDS K22 #341         10,000         2-fam-26         164,474         -         -           AUGERDS K22 #341         10,000         2-fam-26         164,474         -         -           AUGERDS K22 #341         10,000         2-fam-26         164,427         -         -           AUGERDS K22 #344         10,000         2-fam-36         164,427         -         -           VESI 50,242         2-fam         164,612         2-fam<27	Well         TO         CompA Date         GAG         Oil.         WATER PMUL Stormshim         GAG         CompA PMUL Stormshim         WATE PM PMUL Stormshim         GAG         CompA PMUL Stormshim         GAG         CompA	Viell         TO         Comps. Date         GAG         OB.         VIATER bit Port rest bit Port rest bit         GAG         OA3 (b) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c) (c)	Wetl         TO         CompL Date         GAS         Oil.         WATER bit         GAS         GAS	Weilt         TD         Compt. Date         GAS         OIL         PARTER bit         GAS         GAS         GAS         GAS         GAS         PARTER bits rate         GAS         GAS         GAS         PARTER bits rate         GAS         GAS         PARTER bits rate         GAS         GAS         GAS         GAS         GAS         GAS         PARTER bits rate         PARTER bits rate         PARTER bits rate         GAS         GAS         GAS         GAS         GAS         PARTER bits rate         PARTER bits rate         PARTER bits rate         GAS         GAS	Weft         TO         CompL Date         GAS         OIL         WATER (bit)         GAS         GAS         GAS         GAS         GAS         Mar / Bit         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         GAS         GAS         GAS         GAS         GAS         GAS         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         GAS         GAS         GAS         GAS         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         GAS         GAS         GAS         GAS         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         GAS         GAS         GAS         GAS         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         Stray Jane (bit)         GAS         GAS         GAS         GAS         GAS         GAS         Stray Jane (bit)         Stray Jane (bit) <thstray jane<br="">(bit)         Stray Jane</thstray>	Weft         TO         Compl. Date         GAS         GAS <th< td=""><td>Viell         TO         Compt. Date         GA         Oil.         MATER IMATER         GAS bit Mode         GAS (reg rate (mage)         GAS (mage)         GAS (reg rate (mage)         GAS (reg rate (mage</td><td>Weft         TO         CompA Date         GAG         OAL         MATER bit         GAS         GAS         GAS         GAS         GAS         GAS         Bit rate         Off         Off</td><td>Weft         TD         Compl. Date         GAS         <th< td=""></th<></td></th<>	Viell         TO         Compt. Date         GA         Oil.         MATER IMATER         GAS bit Mode         GAS (reg rate (mage)         GAS (mage)         GAS (reg rate (mage)         GAS (reg rate (mage	Weft         TO         CompA Date         GAG         OAL         MATER bit         GAS         GAS         GAS         GAS         GAS         GAS         Bit rate         Off         Off	Weft         TD         Compl. Date         GAS         GAS <th< td=""></th<>

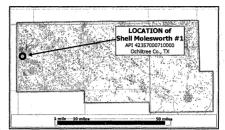
## Atoka Production Statistics:* Vertical - Horizontal Well Comparison

n, # wells	Gas (Bcfg)	Oii (MMBO)		Cum.Gas	Cum.	Cum.	Days			
			(IMIM DAA)	(MMctg)	Oli (MBO)	Wtr (MBW)	on Prod.	Gas (Mcfg)	OII (BOPD)	Wtr [,] (BWPD)
525	219.4	10.80	0.68	417.2	20.6	: <b>2.3</b>	3,045	180.4	13.7	2.3
20	5.0	0.03	0.08	949.9	1.6		382	949.9	4.7	19.1
					I		<u> </u>		<u></u>	20         5.0         0.03         0.08         949.9         1.6         382         949.9         4.7           s: IHS Energy, DrillingInfo ; Lipscomb and Ochiltree Co.'s, TX, Ellis Co., OK; current to Ma

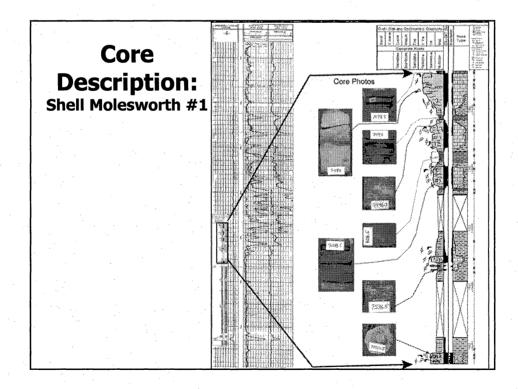
		Ato	ka IP	Stat	istics	*			
V E R		IP GAS IP OIL IP WTR IP BHP mcfgpd bopd bowd psi IP BHT °					Average Perf Depth (MD, ft)		
T I	average	783	45	4	1,626	193	9,282		
C A	min	0	0	0	61	139	7,587		
Ĺ	max	22,000	818	145	4,562	257	13,475		
	aveiage	2,765	28	66	1,335	168	10,638		
	min	22	1	0	389	114	9,811		
	max	7,036	128	360	3,559	255	11,528		
HORIZ	ZONTAL	IP GAS mcfgpd	IP OIL bopd	IP WTR bowd	IP BHP psi	IP BHT °F	Average Perf Depth (MD, ft)		

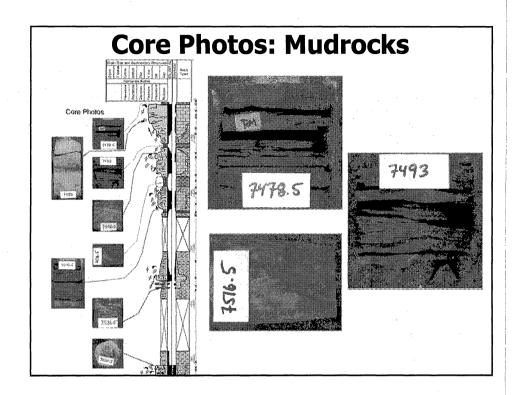


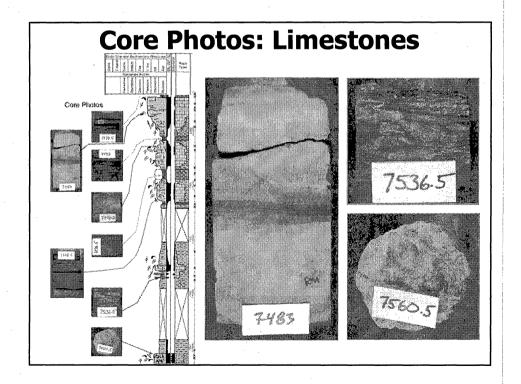


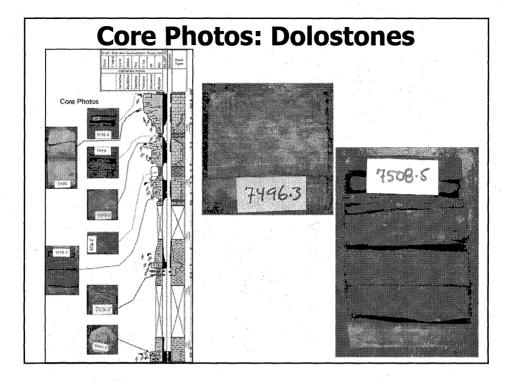


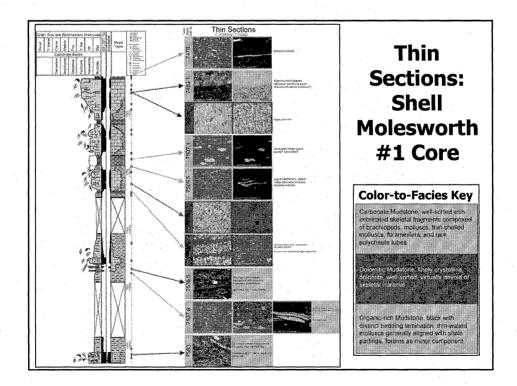
CORE CONDITION.--This dry hole was drilled in 1957 on the western edge of **Ochiltree County. Shell cored 84 feet of** 'Lower Atoka' (informal name used herein) and 786 ft of Morrow. There was fair core recovery in Atoka: 70 ft (83%) was recovered but there are several substantial missing sections. Some recovered sections are rubbly. However, this is the only Atoka core we were able find and based on resistivity log comparison to other wells, the facies in the Molesworth #1 core are probably representative of the Atoka in general.

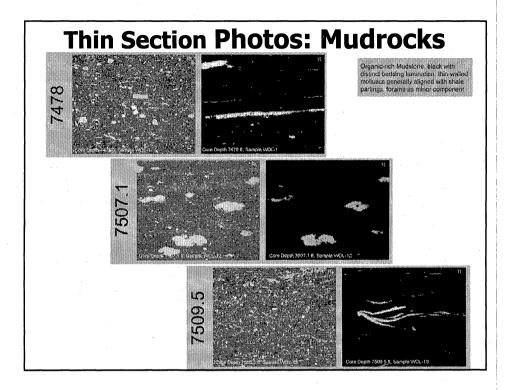


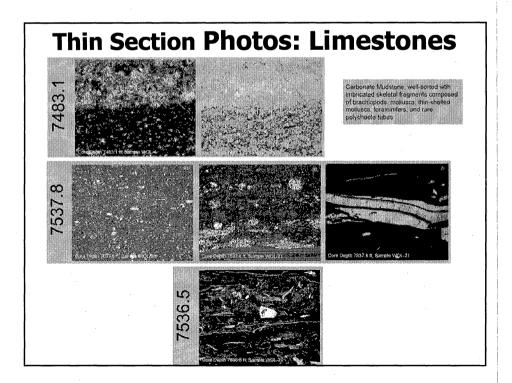


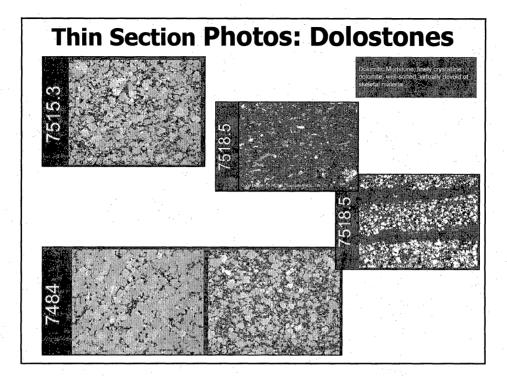












# Interpretation of Depositional environment: Shell Molesworth #1 Core

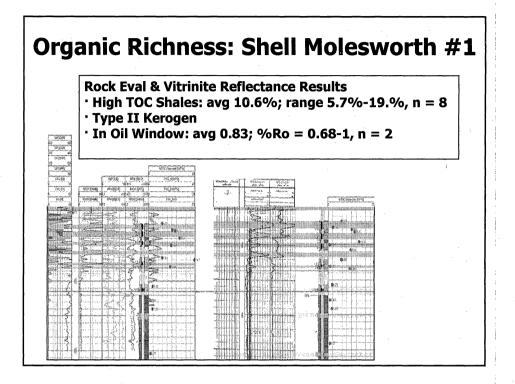
**Deep-water environment:** 

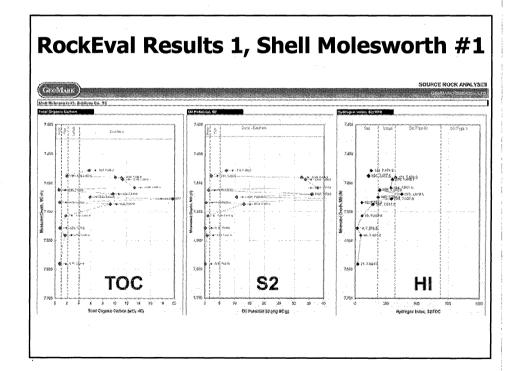
• Restricted, anoxic deep marine basin

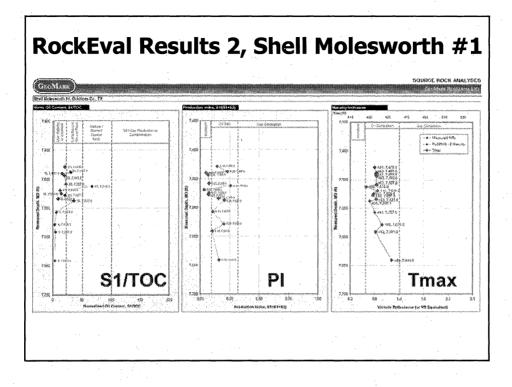
• Interrupted by episodic transport of fine material probably by density flows

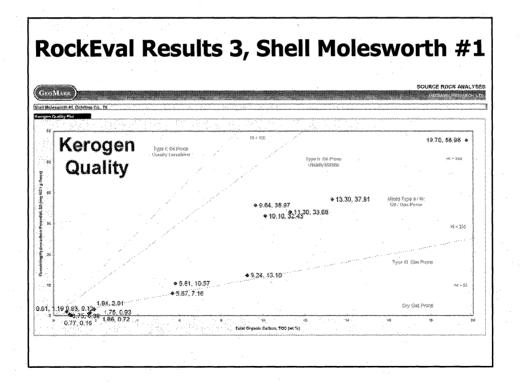
Most transport grains derived from basin-rimming carbonate complexes
Minor siliciclastic input

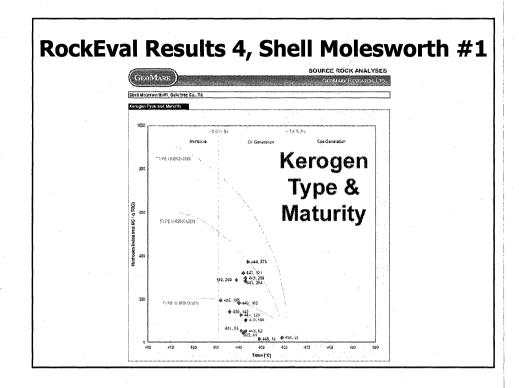
→ Low-order marine condensed section

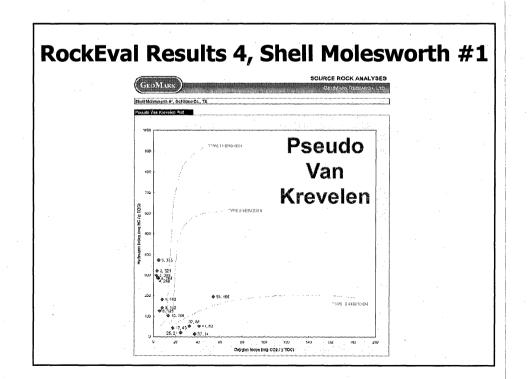


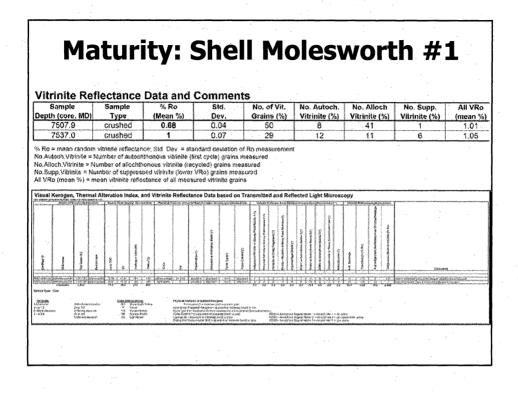




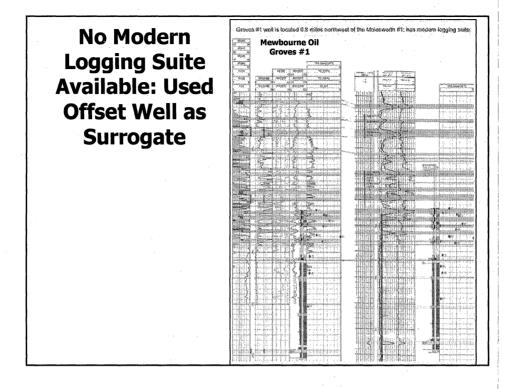


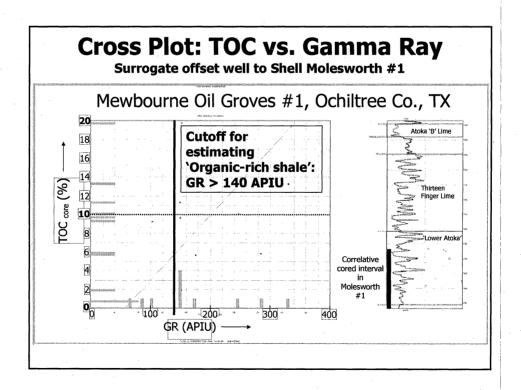


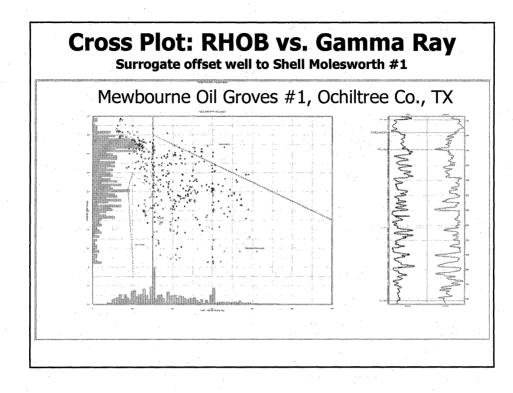


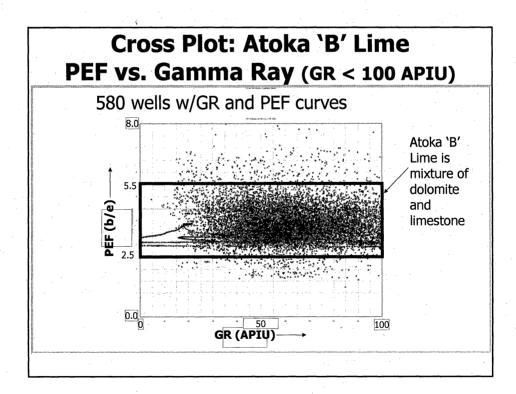


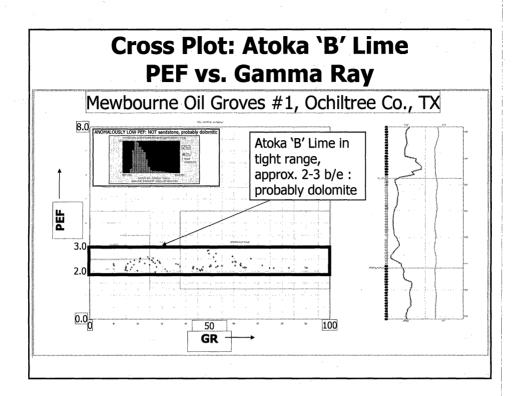
										4			
Gool											SOURCE	ROCK A	NALYSE
Conn											Cost	41.00 P.005	0.977-11 ¹ 1 1
inell Moless	rontli A1. O	amitree Co., TX	·····	·····									
Formation	Depth (ND f6	Lithology OLC intero.1	Laco TOC (wt% HC)	S1 (me HC/g)	Rock-Eval 82 (mg HCa)	S3 (mit CO2/e)	Measured %Ro Warinite Refu		Hydrogen Index (S2x100/TOC)	Oxygen Index (S3x100/TOC)	S2/S3 Conc. Ima HC/ma CO2)	St/TOC Norm. Oil Content	Productio index (S1/(S1+5
1 Atoka		isiank sisala	5.67	1 / 9	7 \$R	0.82		441	125	9	<u>77</u>	21	
L. Atoxa	7,497 5	cal/arsous shale	1.94	187)	2.61	0 28		443	102	(3	Ą	<b>3</b> 7	1
L. Aloka		black shale	10.13	1.62	32.43	0.34		412	321	3	95	10	3
1. Atoxa		heark anala	1: 30	2.28	33.68 37.81	0.31	0.34	443 443	206	3	105	20 23	7 2
L. Aloka		black shale dx cray limestone	0.61	0.51	- 10	0.55	0,94	432	195	54		87	, 3
E. Alexa		plack choice	2.61	· 1.20	35.97	0.48		434	372	5	73	32	э
i. Atoka	7,524.4	hinck shale	5.81	D 84	10.57	0 37		440	165	<b>Ŗ</b>	21	14	э
L Atoka	7,527.3	calcarecus shale	19 75	4.94	58.96	0.78	a anaina anna a	439	380	1	73	21	С
L. Atoxa		dk prey lin estens	0.76	5 67	0,30	0.31		643	52	.41	1	ø	с С
L Atoko		block shole	0.24	3.37 0 (4	13.10 0.93	0 73 0.87	1.00	436	142	8 33	18	96 0	2
L. Aloka		calcarecus chalo	0.83	. 3.63	0.10	0.5/ 0.80		440	59  4	87	۸ ۵	4	2
U. Morew		calkaroous shale	1.66	3.96	0.72	0.20		44.2	4.3	:7	2	6	3
		calcerecus shelp	0.77	2.83	0 16	0,19		489	21	26		4	S

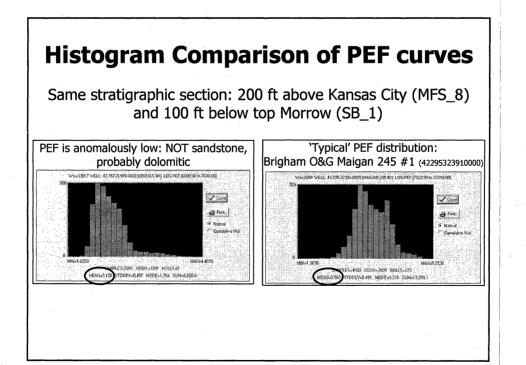


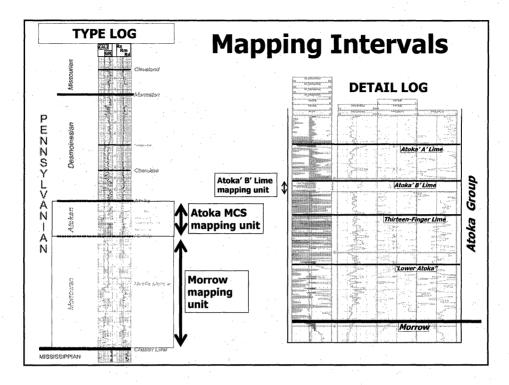


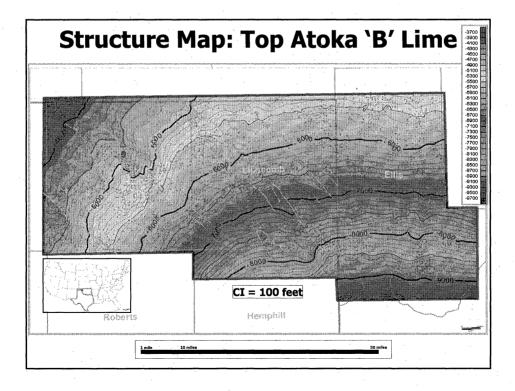


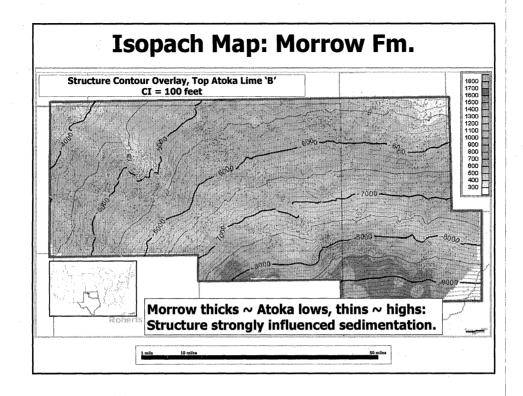


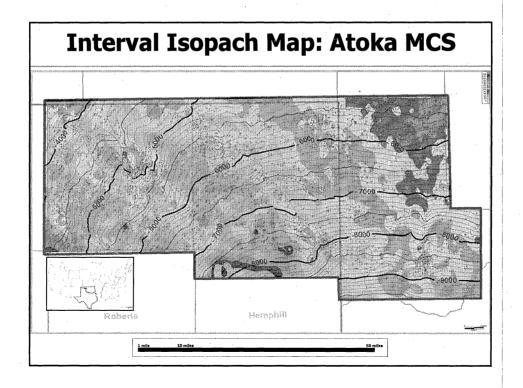


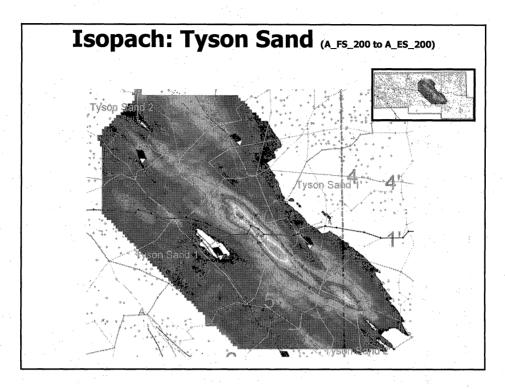


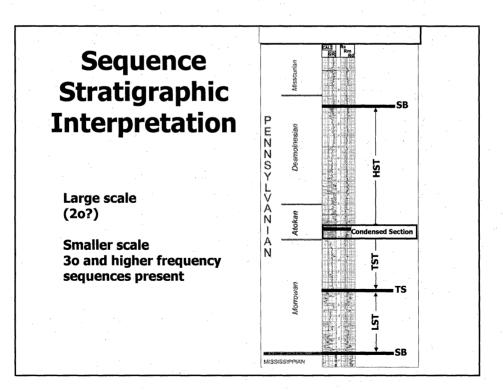












## Summary

Study represents a reconnaissance of Atoka Lime/Shale geology. Need more whole cores to fully assess but, several conclusions can be made from the work thus far:

Activity:

- Historically a minor secondary target
- Vertical-wells in locally-developed porosity in the Atoka 'B' Lime (aka Novi Lime)
   Emerging horizontal/ fracture stimulation has thus far increased average gas production rates 5-fold. Some of this increase may be a result of hydraulic fractures reaching into the adjacent organic-rich shales.

## Mudrocks:

Highly organic-rich, fissile black shales.

- Calcareous and fossiliferous adjacent to carbonate beds
- Shell Molesworth #1 core
  - TOC avg 10.6%
  - Vitrinite reflectance averages 0.84 (oil window)
  - Average measured depth of 7522 ft
  - Estimate of total organic-rich shale footage in the Atoka done via 1100 wells with LAS-GR
  - data yielded an average of 214 ft of 'hot shale' (GR > 140 APIU)
  - 'Upper Atoka' contains the most 'hot shale', averaging 56 feet

Carbonates:

• Shell Molesworth #1 'Lower Atoka' core: carbonates are deep-water limestones were formed by transport of carbonate skeletal debris

- •Very low porosities even though many were highly dolomitized
- No core-no direct observations of Atoka 'B' Lime but have observed that:
  - •Averages: isopach 11 ft thick, up to 37 ft; and averages 7% (limestone matrix) porosity •Highest porosities, up to 25%, are typically found in dolomitized zones
    - •Thickest carbonate isopach values in present-day structural lows
      - •Structure controlled local accommodation in Atoka and underlying Morrow Fm.

## Summary (cont.)

• Depositional Environments:

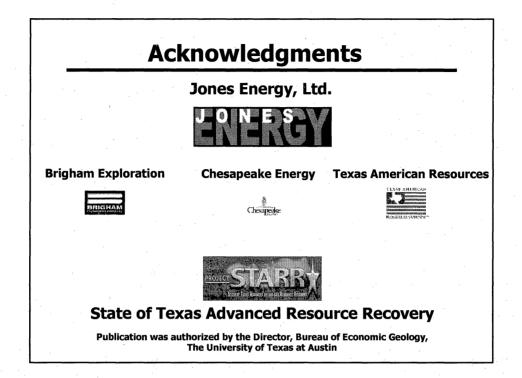
• Deep-water basinal position at the northwestern terminus of Anadarko Basin

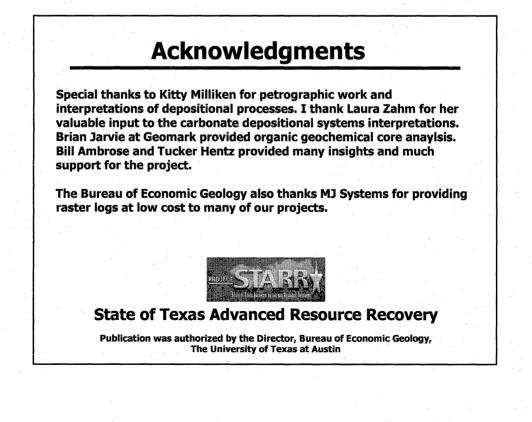
anoxic, restricted deep-basin shale deposition was interrupted by frequent pulses of carbonate
 sedimentation and volumetrically minor thin sandstones, via turbidites

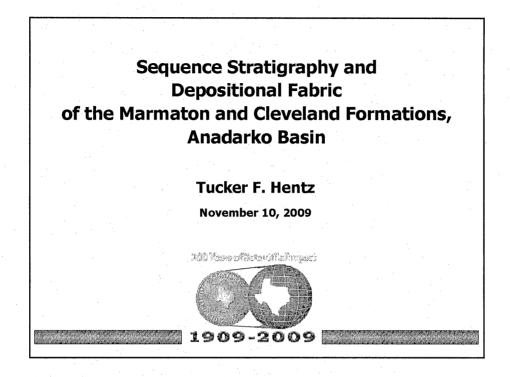
• Carbonate sediment sources were probably from adjacent shallow shelf/platforms along the north and north-eastern margins of the Anadarko Basin.

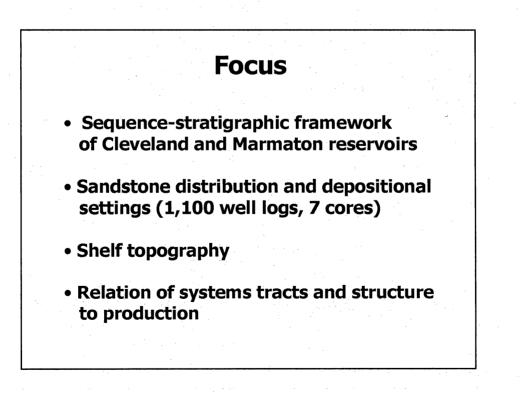
Sequence Stratigraphy:

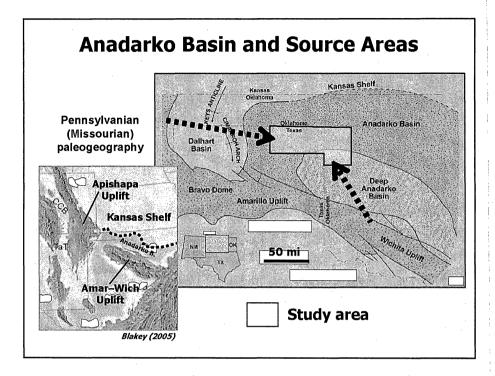
- Atoka Fm. was deposited during a major global transgressive during the Middle Pennsylvanian
   Represents a large order (20?) marine condensed section
- Many higher frequency cycles
- Sequences are volumetrically dominated by HST deposits

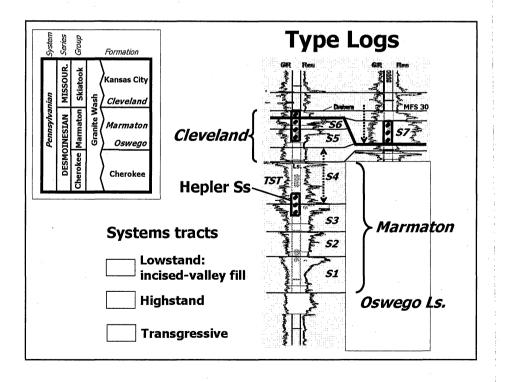


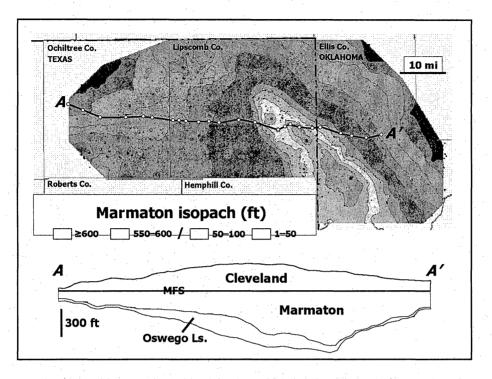


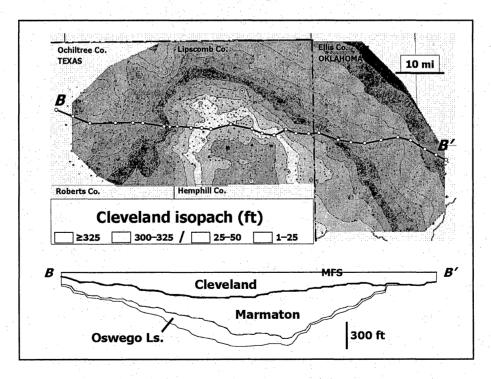


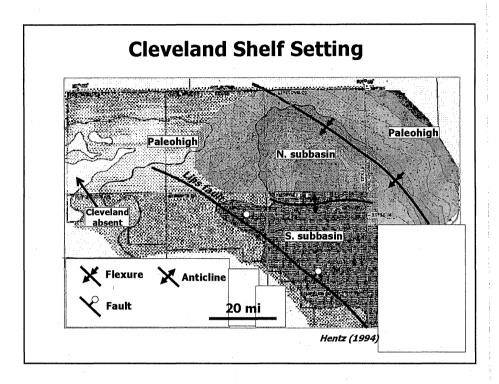


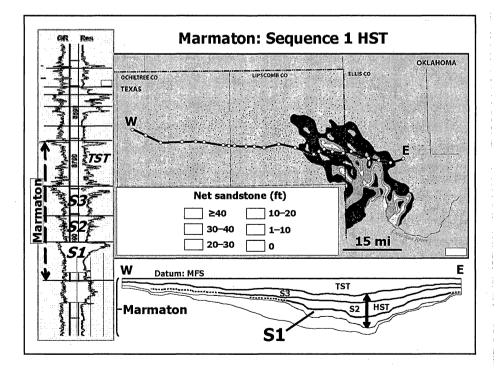


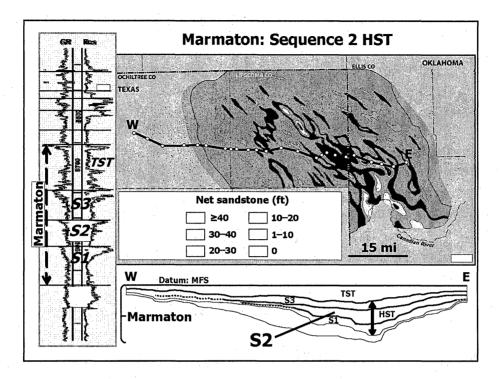


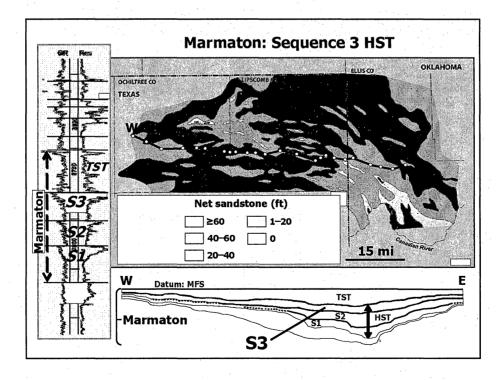


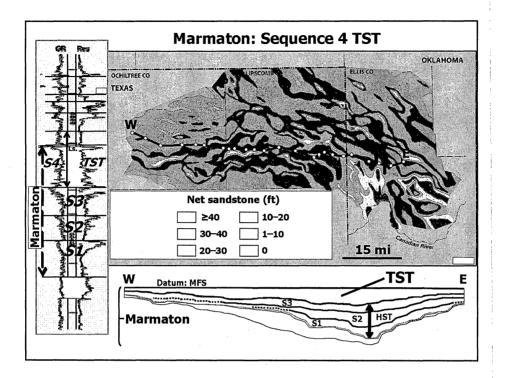


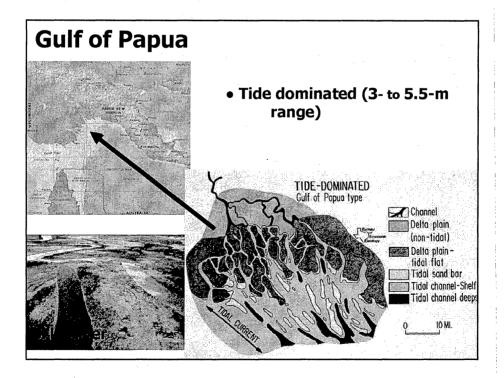


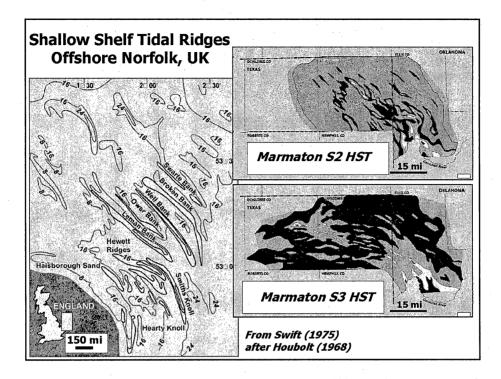


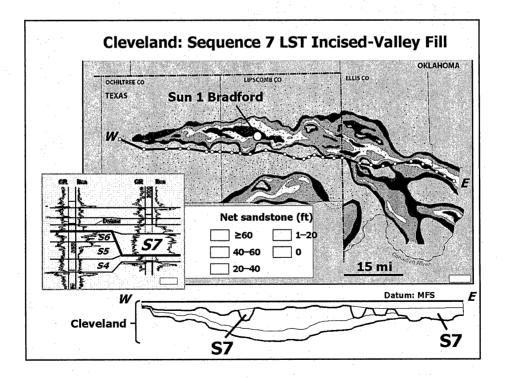


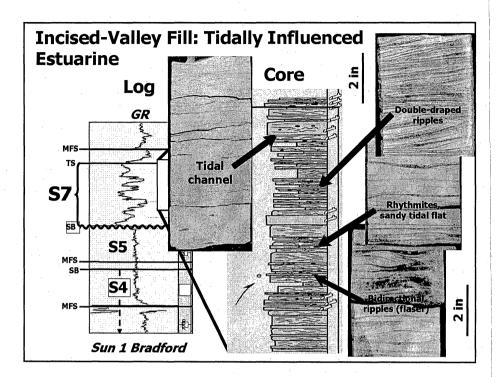


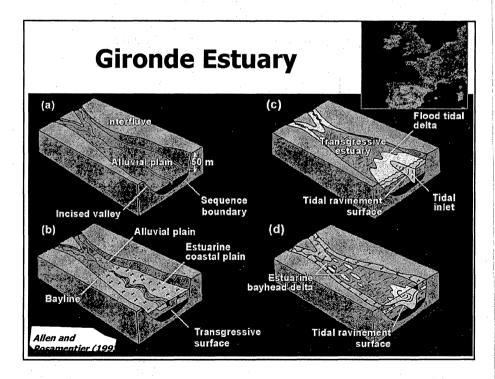


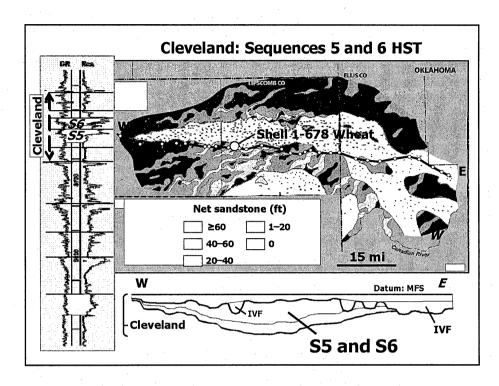


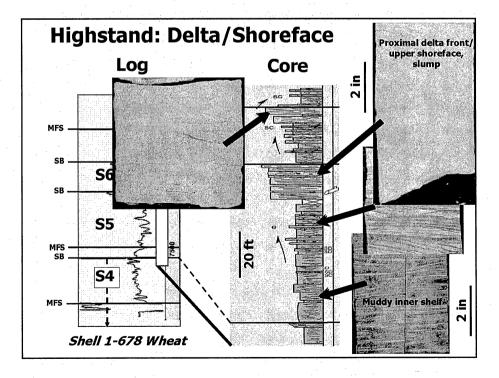


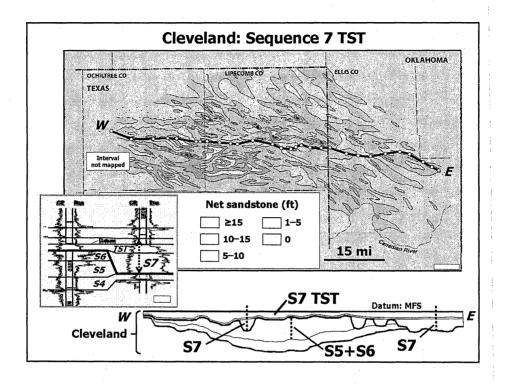


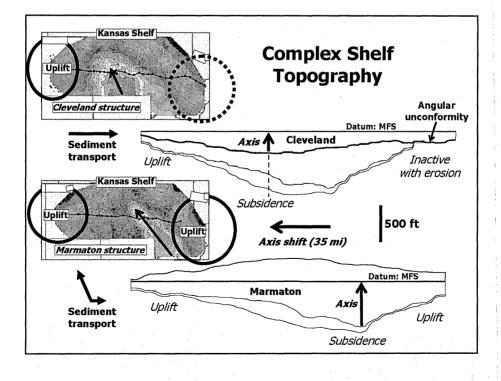


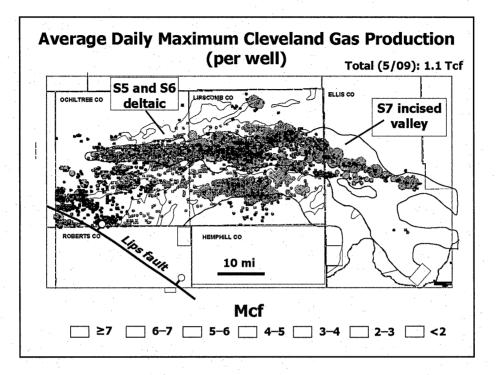


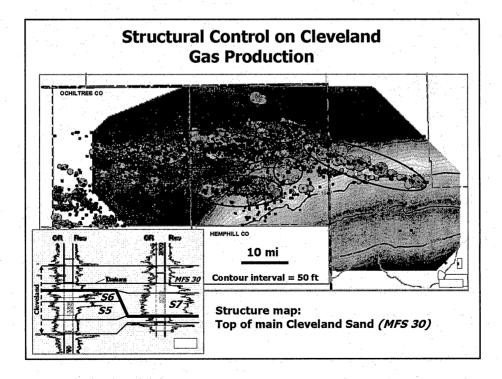


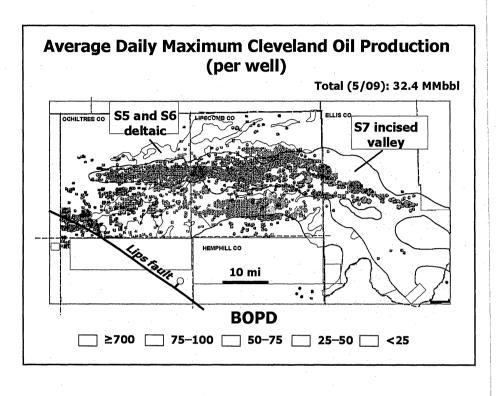


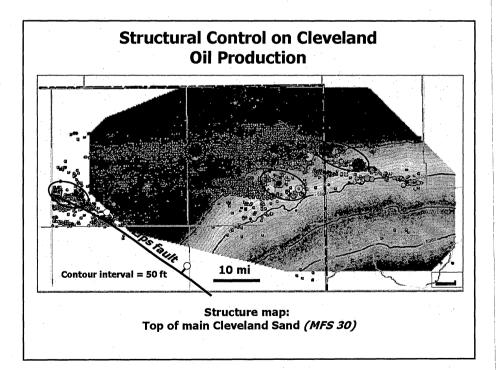


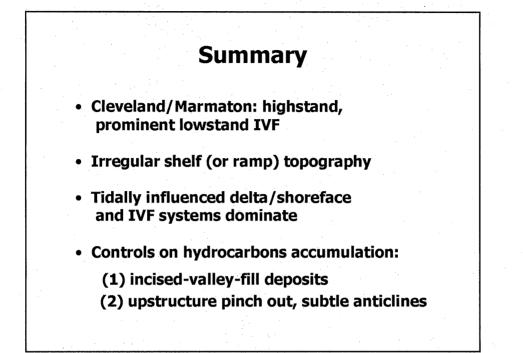


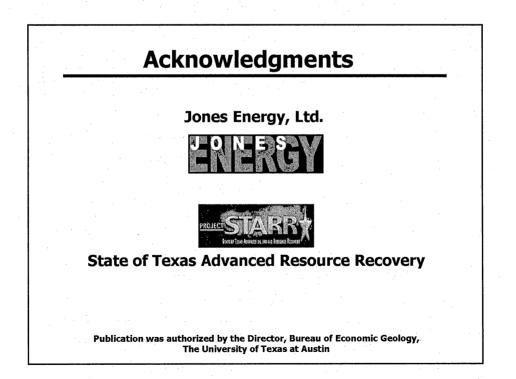


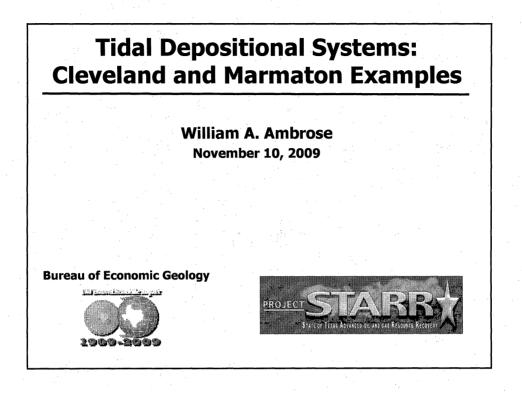


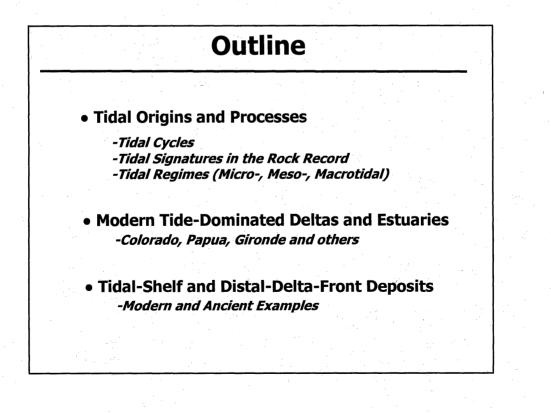


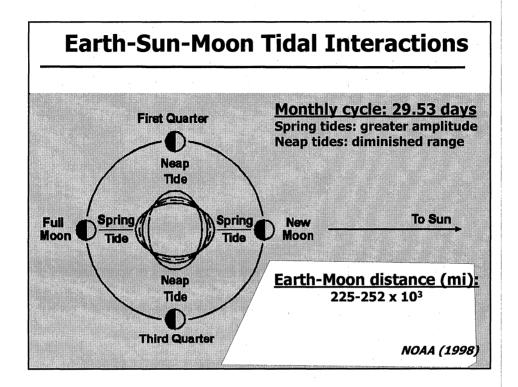


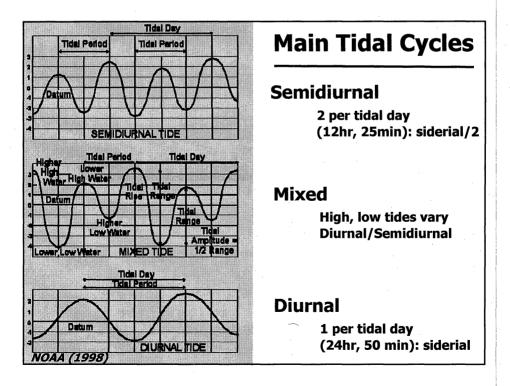


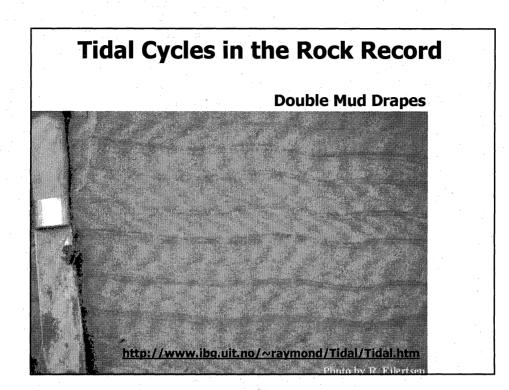


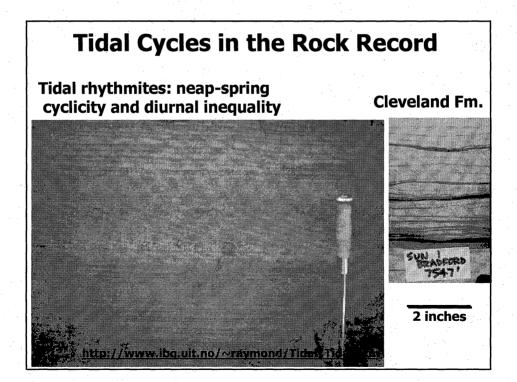


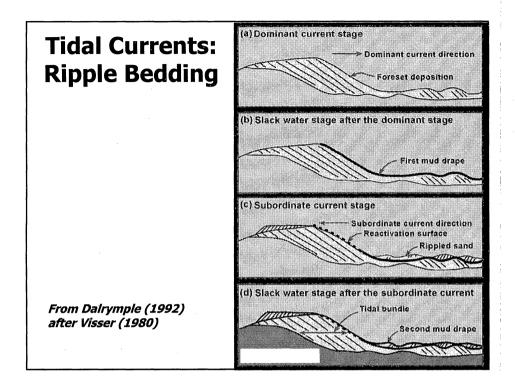


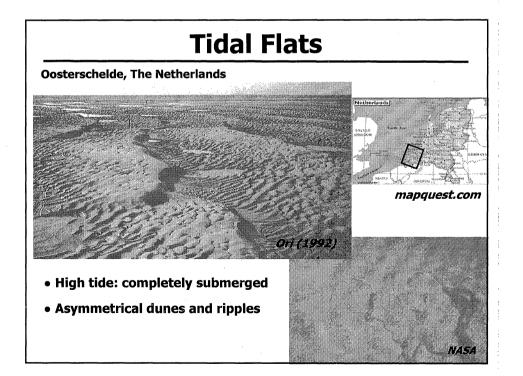


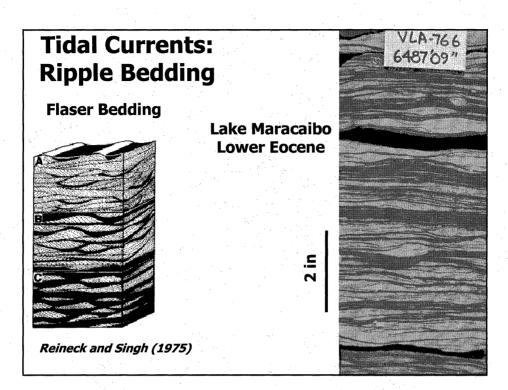


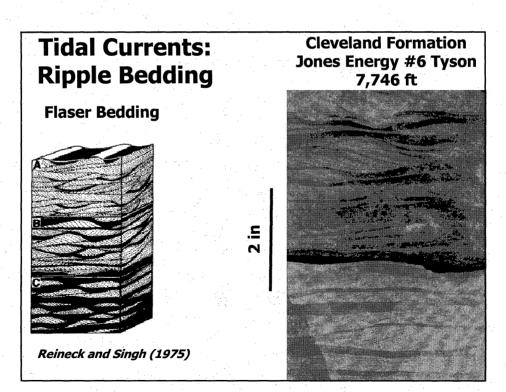


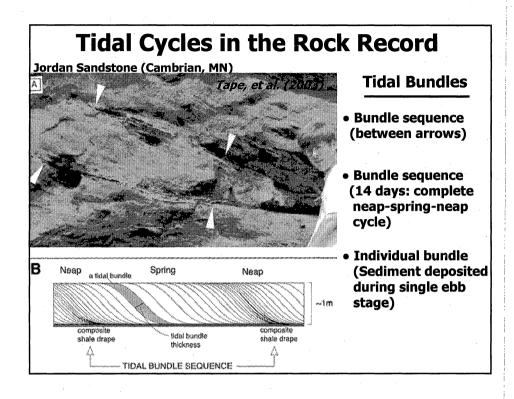


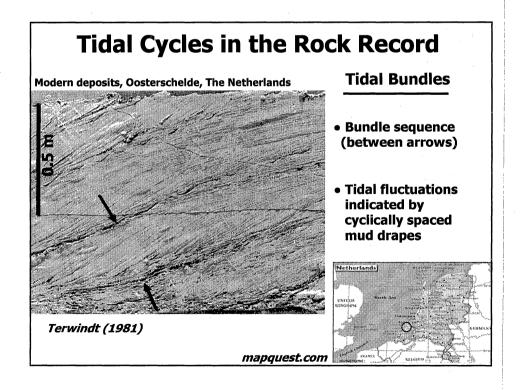


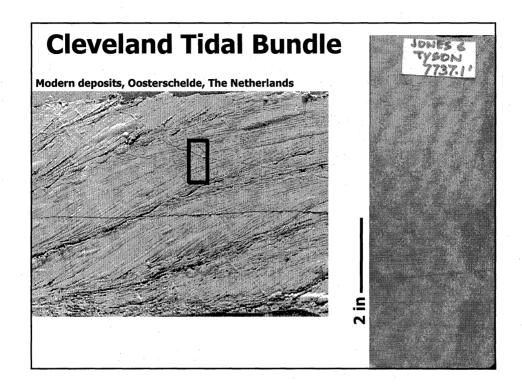


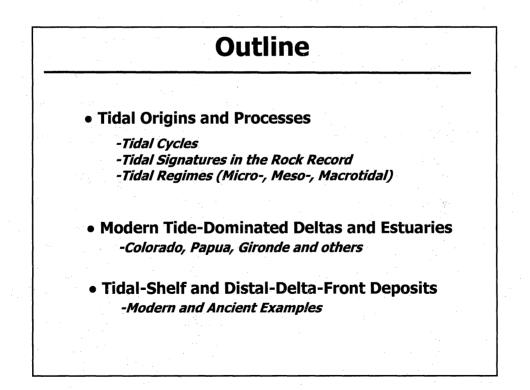


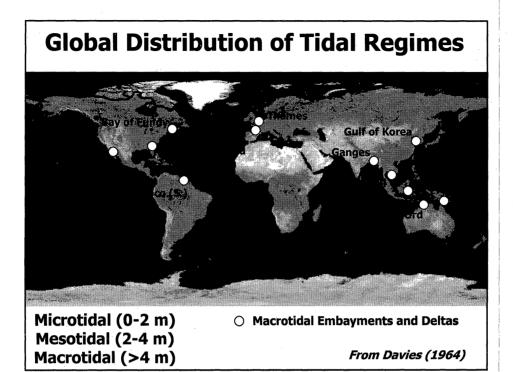


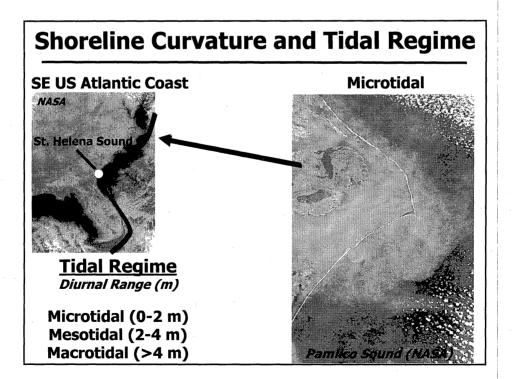


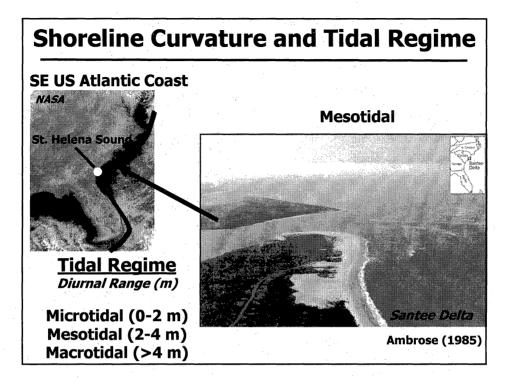


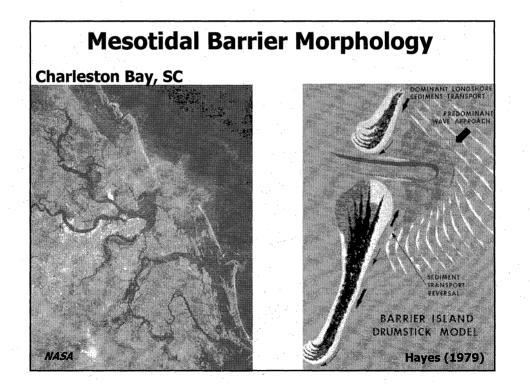


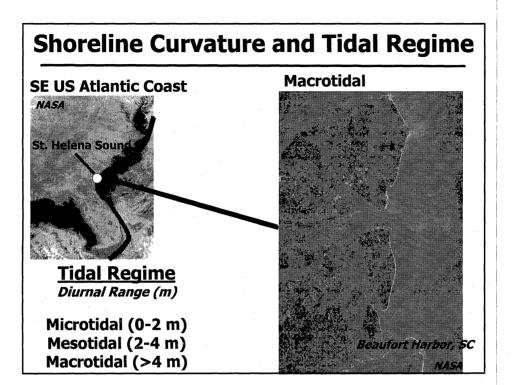


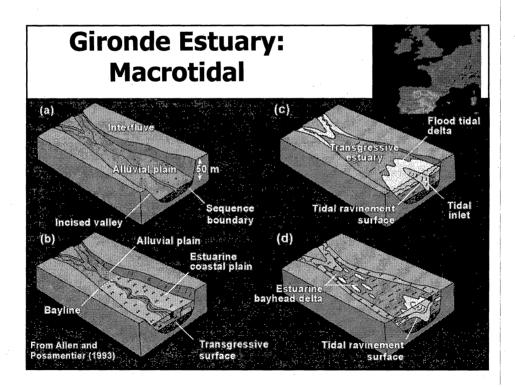


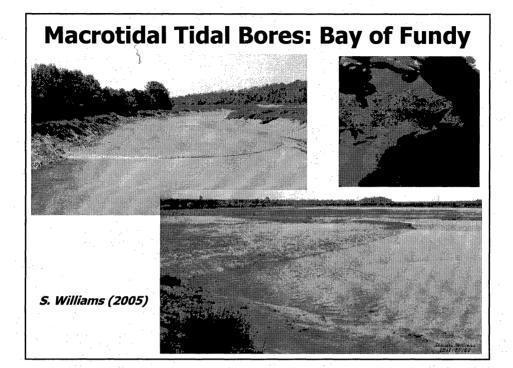


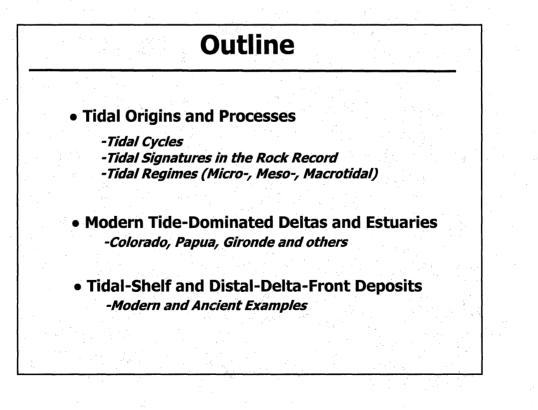


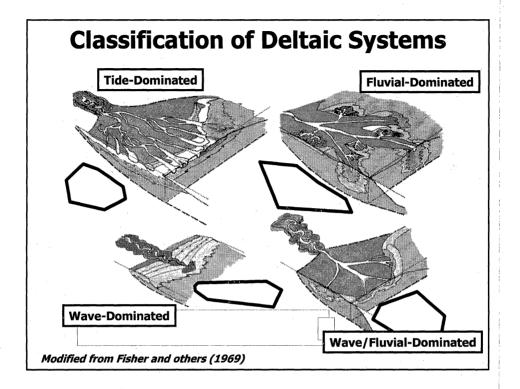


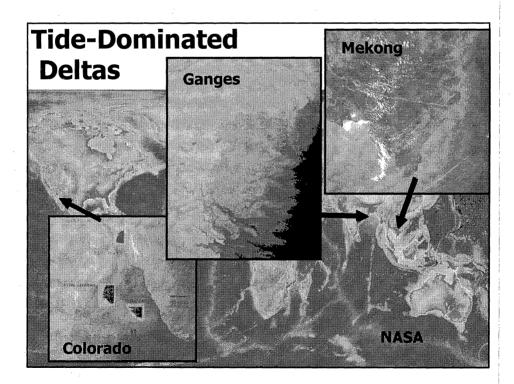


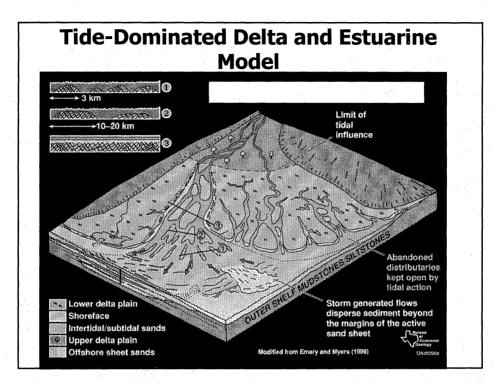


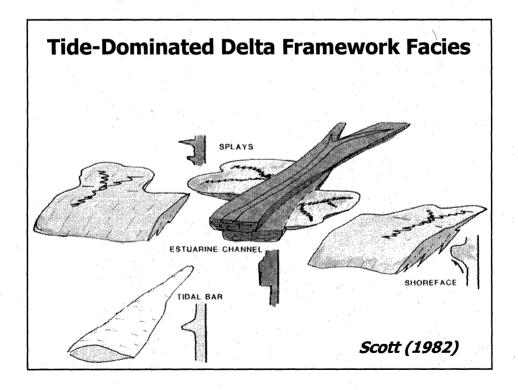


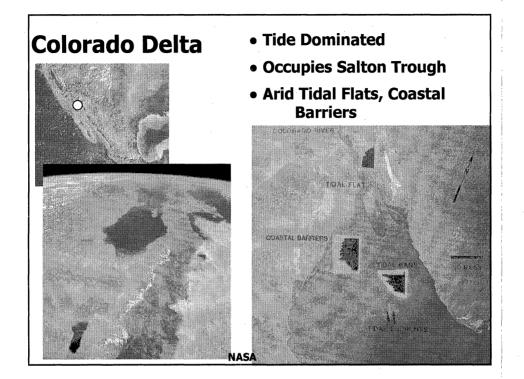


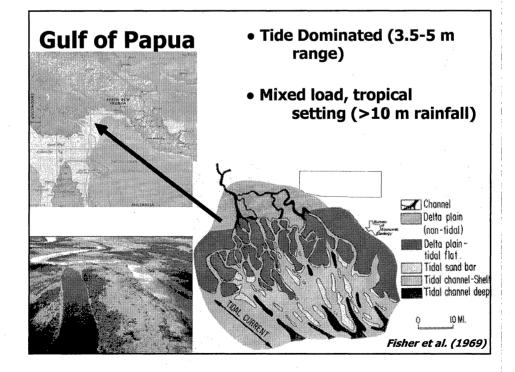


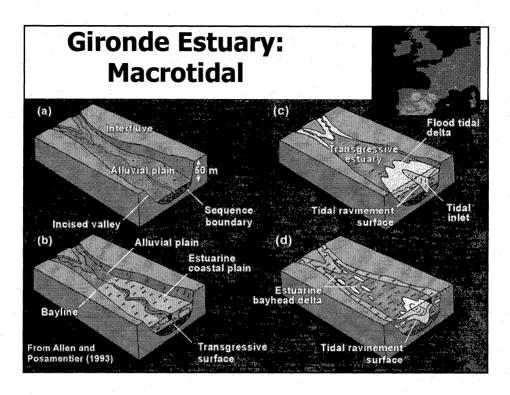


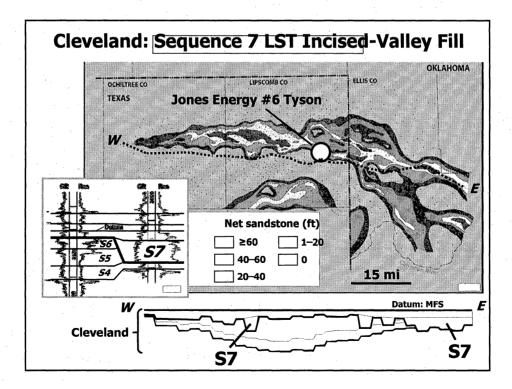


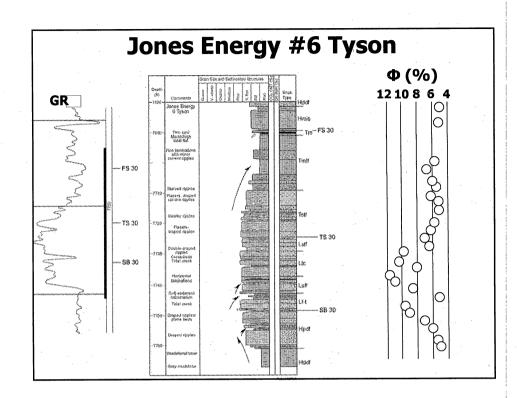


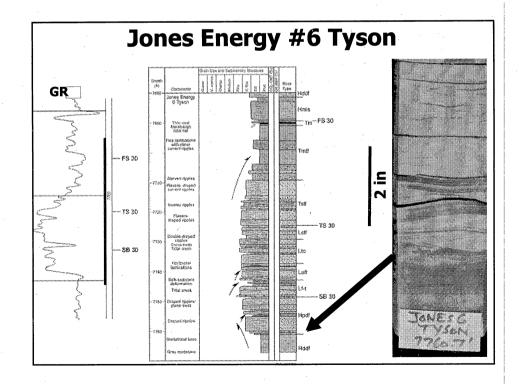


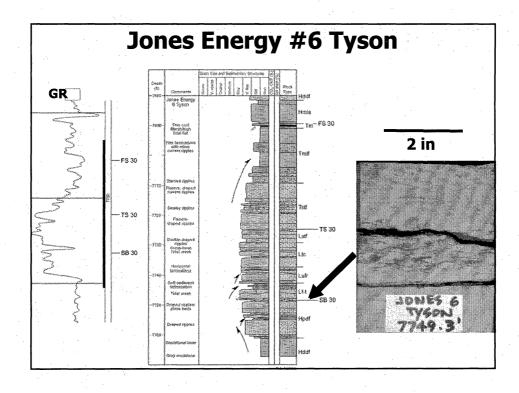


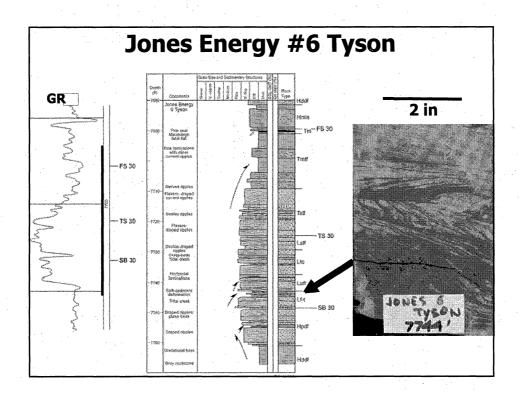


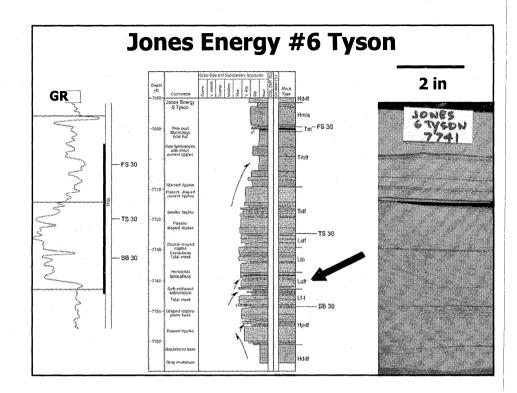


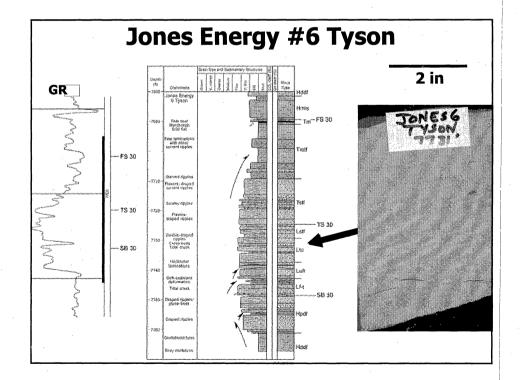


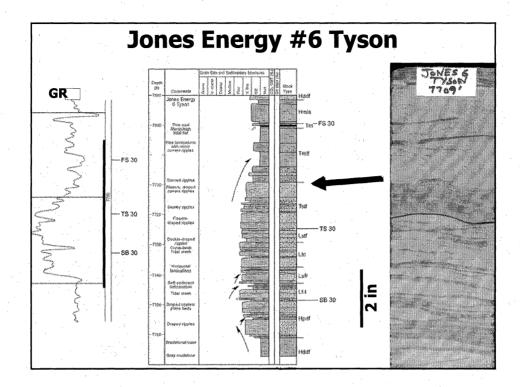


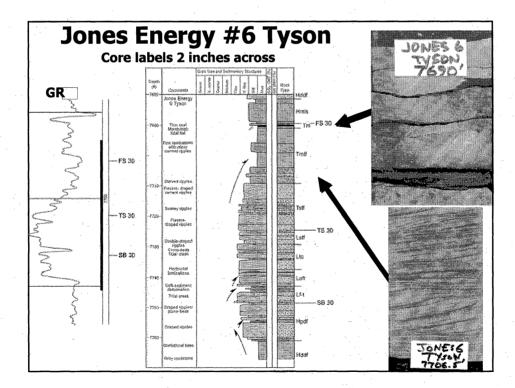


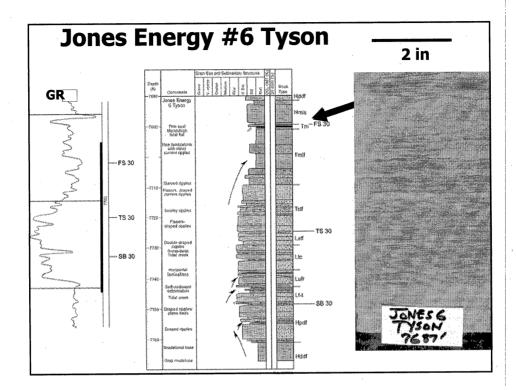


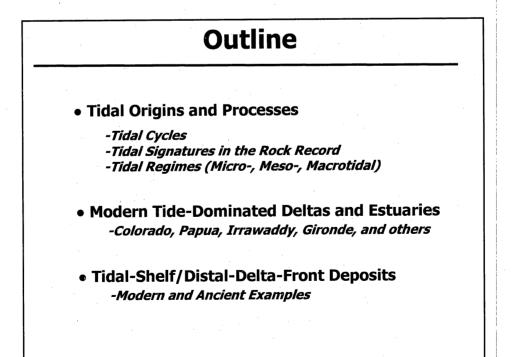


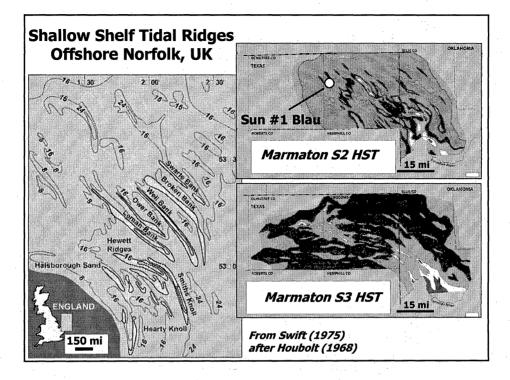


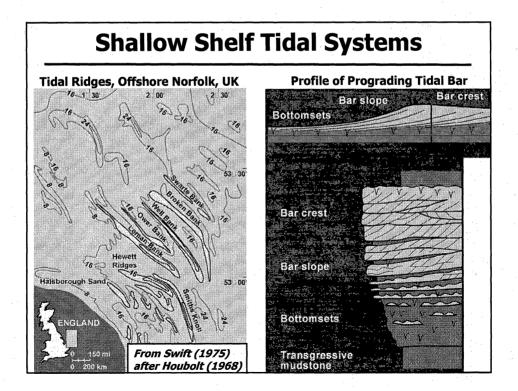


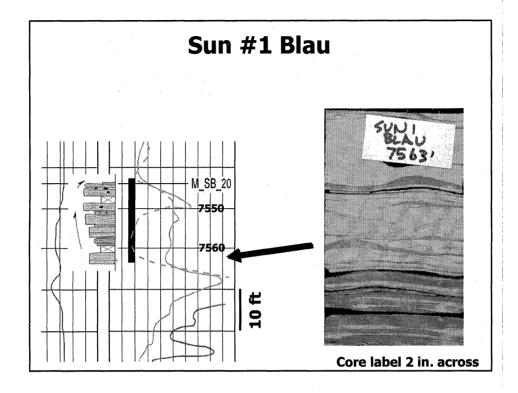


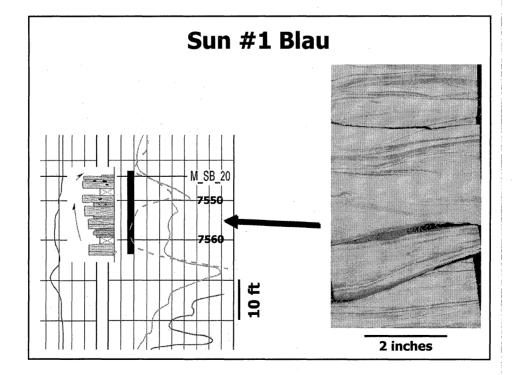


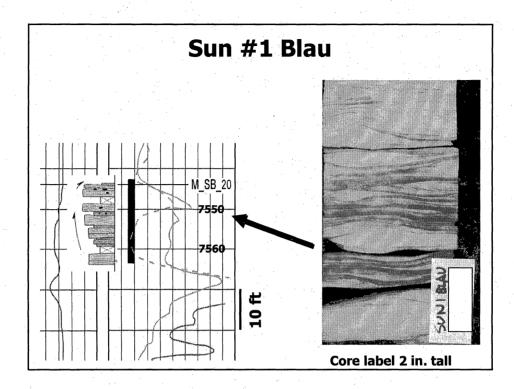


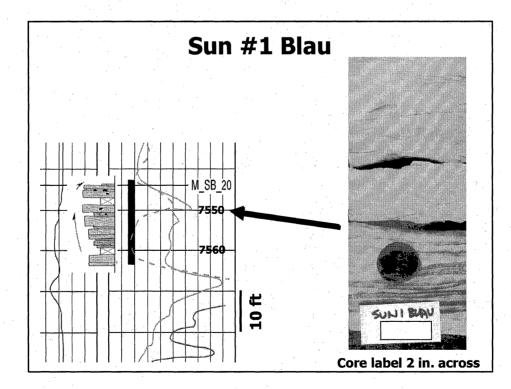












## Summary

- Tidal Origins and Processes
  - -Tidal Cycles
  - -Tidal Signatures in the Rock Record
  - -Tidal Regimes (Micro-, Meso-, Macrotidal)
- Modern Tide-Dominated Deltas and Estuaries -Colorado, Papua, Gironde and others
- Tidal-Shelf and Distal-Delta-Front Deposits -Modern and Ancient Examples

### Core Workshop Cleveland Formation and Marmaton Group Anadarko Basin

#### William A. Ambrose and Tucker F. Hentz

Bureau of Economic Geology John A. and Katherine G. Jackson School of Geosciences The University of Texas at Austin Austin, TX

November 10, 2009

#### Introduction

Cores in the Cleveland Formation and Marmaton Group in the Anadarko Basin display a variety of facies in tidally dominated depositional settings. This workshop presents cores that illustrate contrasting sequence tracts in the Cleveland Formation that include the Jones Energy #6 Tyson (lowstand-valley-fill and transgressive-estuarine deposits) and the Shell #1-678 Wheat (tide- and wave-modified, highstand-shelf, and delta-front facies). In addition, a short cored section from the Sun #1 Blau well features inner-shelf and delta-front highstand deposits from the Marmaton Group (Fig. 1). Another core from the Atoka Series, the Shell #1 Molesworth in Ochiltree County (located in Fig. 1) demonstrates variable carbonate and shale lithology and is described in a separate document in this guidebook.

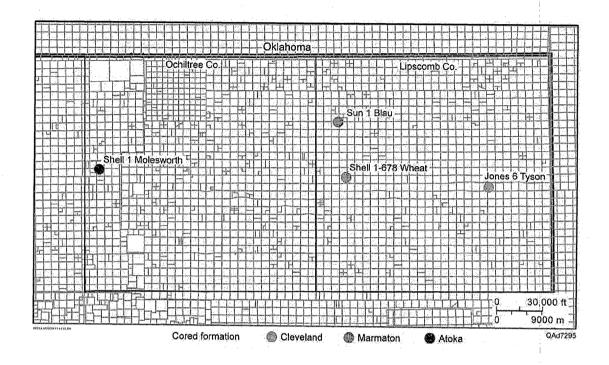


Figure 1. Location of cored wells presented in this workshop.

Cores are vital to an understanding of reservoir architecture because they display a vertical profile of the rock fabric, contacts between beds, and sedimentary structures from which one can infer sedimentary processes and depositional facies. Vertical trends in lithology, grain size, and stratification are the result of changes in depositional environments through time. Workshop participants will have the opportunity to observe numerous examples of vertically superimposed sedimentary facies in the Cleveland Formation and Marmaton Group in a variety of depositional settings that include incised valley, estuary, delta, and highstand shelf.

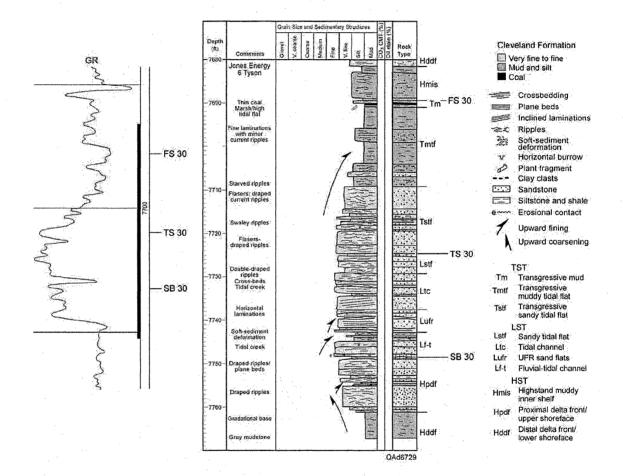
Recognition criteria for tidal deposits in the Cleveland Formation and Marmaton Group include distinctive stratification and bedforms observed in core, as well as sandstone-body geometry inferred in lithofacies maps of distal-deltaic deposits in the lower Cleveland highstand sequence and shallow-marine deposits in the upper Cleveland transgressive sequence, where the geometry of these sandstone bodies reflects the relative strength of tidal- versus wave-reworking processes. Stratification and bedforms diagnostic of tidal deposits in Cleveland cores include asymmetric, double-draped ripples with bidirectional foresets; abundant reactivation surfaces; rhythmic, laminar stratification; and flaser bedding. These types of features have been described in other studies of both modern and ancient tidally influenced deposits, including Reineck and Wunderlich (1968), Klein (1970), de Mowbray and Visser (1984), Dalrymple and others (1990; 1991), and Allen (1991). Asymmetric current ripples in cores of the Cleveland Formation in particular, interpreted to be tidal in origin, typically exhibit differences in degree of development and size between ripple sets of alternating orientation. These differences in magnitude of the reversal of bedforms reflect the presence of dominant and subordinate tidal currents, which are common in many estuarine settings as a result of the relative strength of flood- and ebbdominated bedform migration in different reaches of the estuary. For example, in upstream estuarine settings such as in the Ord River Delta in Australia, the velocity of flood-tidal currents commonly exceeds that of ebb-tidal currents, resulting in a weak bidirectional ripple fabric (Wright and others, 1973; 1976). In contrast, at the estuary mouth, flood and ebb currents are approximately equal in strength, and a strong bidirectional ripple orientation is therefore more clearly evident (Hayes, 1976).

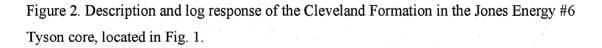
Other features in cores of the middle Cleveland lowstand, incised-valley-fill sequence that are consistent with a tidally influenced depositional setting include abundant sections of plane-bedded, laminated sandstone interpreted to represent upper-flow-regime (UFR), sandy tidal-flat deposits (UFR sand-flat facies), similar to that described by Dalrymple and others (1992) and Allen and Posamentier (1993) in the modern Gironde Estuary. Although the presence of plane-bedded, laminated sandstone is not in itself indicative of tidal conditions, its association with strata composed of flaser ripples, double-draped ripples with reactivation surfaces, and rhythmic, laminated bedding is consistent with tidally influenced deposits. Additional evidence of tidally influenced depositional systems is from sandstone-body geometry in lithofacies maps of distal-delta-front and shelf deposits in the lower Cleveland highstand sequence, composed of upward-coarsening, progradational successions exhibiting dip-elongate and subparallel net-sandstone-thickness patterns, consistent with reworking of sandy, delta-front sediments from tidal currents in a shallow-shelf depositional setting. Modern examples include elongate, subaqueous tidal bars in the North Sea (Houbolt, 1968; Swift, 1975), the modern Gulf of Korea (Off, 1963), the Gulf of Papua (Fisher and others, 1969), the Ord River Delta in Australia (Wright and others, 1973), and the Gulf of California (Meckel, 1975). Similar net-sandstone geometries in tide-dominated delta-front depositional settings were documented by Maguregui and Tyler (1991) and Ambrose and others (1995) in Eocene tide-dominated deltas in Lake Maracaibo, Venezuela, as well as tide-dominated shelf and deltaic deposits in the Eocene Baronia deltaic system in northern Spain (Mutti and others, 1985) and Eocene tidal-sand-ridge deposits in Belgium (Houthuys and Gullentops, 1988).

4

#### **Cleveland Formation: Jones Energy #6 Tyson Core**

The Jones Energy #6 Tyson core (located in Fig. 1) offers an excellent example of a continuous, 87-ft (~23-m) section through a tidally influenced, incised-valley-fill system (Fig. 2). The basal 17 ft (5.2 m) of the core consists of upward-coarsening, highstand delta-front deposits inferred to be truncated by a lowstand, incised-valley-fill section that extends from 7,724 to 7,748 ft (2,354 to 2,362 m). This lowstand section represents formation of a paleovalley that was filled with estuarine deposits during subsequent sea-level rise, recorded by an upward-fining section of transgressive deposits that extend from 7,748 to  $\sim$ 7,690 ft (2,362 to 2,344 m). The uppermost 10 ft (3 m) of core is interpreted to be innershelf and distal-delta-front deposits in another highstand sequence above a flooding surface (FS 30) that caps the transgressive section at 7,690 ft (2,344 m).





#### **Highstand Systems Tract**

The basal 17 ft (5.2 m) of the Jones Energy #6 Tyson core is composed of highstand deposits below the base of the lowstand systems tract that is identified by the SB_30 marker. This section is upward coarsening, ranging from weakly laminated, silty mudstone at the base to fine-grained sandstone at the top. Sedimentary structures in this lower highstand succession are interpreted to be wave- and tidal-reworked delta-front deposits on the basis of net-sandstone distribution maps that indicate dip-elongate sandstone body geometry (Hentz and others, 2009) and the presence of both wavy and asymmetrical, draped ripples. Thin (0.5-inch [~1.3-cm]) zones of plane beds also occur toward the top of the highstand section, possibly recording a transition to upper flow regime and higher-energy conditions. The gamma-ray (GR) response of the highstand systems tract is commonly serrate to upward coarsening, reflecting a progradational origin.

#### **Lowstand Systems Tract**

The Cleveland lowstand systems tract, which extends from 7,724 to 7,748 ft (2,354 to 2,362 m) in the Jones Energy #6 Tyson core, is a slightly upward fining section of finegrained sandstone, with thin (commonly <3-inch [ $\sim$ 7.5-cm] beds of siltstone and silty mudstones. The base of the lowstand systems tract is marked by sharp-based, upper finegrained sandstones that fine upward into siltstones with flaser ripples. The lowstand section is a heterogeneous succession of facies composed from bottom to top of fluvial-tidal (Lf-t), UFR (Lufr), tidal-channel (Ltc), and sandy-tidal-flat (Lstf) facies (Fig. 2).

The Lf-t facies, 6 ft (1.8 m) thick, consists of an upward-fining section of erosion-based, upper fine-grained sandstone with planar stratification and asymmetric mudstone-draped flaser ripples inferred to have been deposited by alternating tidal currents (for example, at 7,746 ft). This basal lowstand section is interpreted to represent tidally reworked fluvial deposits at the base of the incised-valley fill. Analogous Lf-t facies in modern estuarine settings include alluvial-channel deposits composed of gravelly sandstones in the Gironde Estuary (Dalrymple and others, 1992; Allen and Posamentier, 1993); gravelly and shelly

sandstones in a channel-fill succession in Ossabow Sound, Georgia (Greer, 1975); and upward-fining, scour-based sandstones in the Cobequid Bay-Salmon River Estuary in the Bay of Fundy (Dalrymple and others, 1990). Alluvial-channel deposits in the basal estuarine fill section in the Gironde Estuary are typically much thicker and compose a greater percentage of the overall estuarine succession than the equivalent Lf-t facies in cores of the Cleveland Formation (Allen and Posamentier, 1993). However, the lowstand alluvial-channel facies occurs at the mouth of the Gironde Estuary, with relatively little or no fluvial aggradation in the upstream part of the estuary. This scenario is consistent with the poorly developed basal fluvial-tidal section in the Jones Energy #6 Tyson core, which occupies an updip part of the overall incised-valley system in the study area (see slide set "Sequence Stratigraphy and Depositional Summary of the Marmaton and Cleveland Formations, Anadarko Basin," this volume).

The section of Lft-facies is overlain by the UFR facies, a 6-ft (1.8-m) section of finegrained sandstone with planar stratification, with particularly good examples from 7,739 to 7,741 ft (2,359 to 2,359 m). The UFR facies is interpreted to represent shallow (typically <6.5-ft [<2-m] water depths), upper-flow-regime tidal currents exceeding 6.5 ft s⁻¹ (2 m s⁻¹) that are associated with fluctuating water levels (Reineck and Singh, 1973; Dalrymple and others, 1990). Parallel-laminated intervals in tidally influenced depositional settings commonly form during maximum tidal flow velocities that occur either at the peaks of flood or ebb tides. In contrast, cross-stratified tidal sediments are commonly developed during accelerating or waning-flow conditions within the tidal cycle (Kreisa and Miolola, 1986). UFR facies are documented in a variety of modern macrotidal estuaries, where the diurnal tidal range is at least 13 ft (4 m) (Davies, 1964). Some examples include the Cobequid Bay-Salmon River estuary in Canada (Dalrymple and others, 1990; 1991) and the Gironde Estuary in southwest France (Allen, 1991). In estuarine systems the UFR facies commonly occurs in the transitional area between the distal tidal-sand-bar facies and the proximal tidal-meanderbelt facies along the axial part of estuaries (Lambiase, 1980; Dalrymple, 1992). The UFR facies can either underlie or overlie the tidal-meander facies. The UFR facies is also recognized in ancient successions of tide-dominated estuarine

deposits such as the Eocene Central Basin in Spitsbergen, where individual deposits consist of river-dominated, 3- to 4-ft (0.9- to 1.2-m) sections of sandstone interbedded with bidirectional, ripple-laminated sandstones in abandoned-channel sequences (Plint-Björklund, 2005).

The overlying tidal-channel (Ltc) facies in the Jones Energy #6 Tyson core is an upwardfining section of fine-grained sandstone with crossbeds, inclined stratification, climbing ripples, clay clasts, and organic fragments. Diagnostic features of the Ltc facies include inclined stratification (for example, at 7,737 ft [2,358 m]), crossbeds (7,731 ft [2,356 m]), and numerous internal scour surfaces, with minor planar stratification overlain by ripples with mud drapes. Tidal-channel successions in modern examples of estuarine systems or areas with significant tidal influence, such as the Ossabow Sound in Georgia (Greer, 1975), the Solway Firth in Scotland (Bridges and Leeder, 1976), tidal flats in the Netherlands in the North Sea (van Stratten, 1954; Reineck, 1967), and the Gironde Estuary (Dalrymple and others, 1992; Allen and Posamentier, 1993) display vertical successions similar to those in the Cleveland Formation, with some minor differences. Tidal-channel deposits in the Ossabow Sound are dominated by basal sections of steeply dipping foresets with mud pebbles overlain by muddy, burrowed, and fine-grained sandstone. In contrast, those in the Solway Firth are composed of basal sections of coarse-grained sandstone with mud and shell clasts. Tidal-channel deposits associated with tidal flats in the Netherlands, commonly highly meandering, exhibit characteristics of point bars, with vertical successions consisting of medium- to coarse-grained sandstone with a basal lag, overlain by interbedded, fine-grained sandstone and siltstone with lateral accretion surfaces. These tidal-channel deposits commonly pinch out into and are overlain by sandy tidal-flat deposits (Reineck, 1967). The sandy tidal-flat facies (Lstf), which constitutes the top of the lowstand sequence in the Jones Energy #6 Tyson core, is composed of a 4-ft (1.2-m) section of fine-grained sandstone. The Lstf facies is typically dominated by climbing ripples with thin mudstone drapes, reflecting migration of sandy sediment. The Lstf facies represents deposition from lower-flow-regime tidal currents, with slackwater suspension sedimentation recorded as siltstone drapes over ripples. Most of these ripples are

asymmetric, reflecting deposition from low-energy, alternating tidal currents rather than from wave processes.

#### **Transgressive Systems Tract**

The upper Cleveland transgressive systems tract (TST) records flooding of the LST estuarine valley-fill system. Numerous examples of transgressive estuarine successions are in the ancient rock record, including the Upper Cretaceous Point Lookout Sandstone in northwestern New Mexico (Devine, 1991); the Albian Paddy Member of the Peace River Formation and lower Shaftesbury Formation in Alberta (Leckie and Singh, 1991); the Lower Triassic in Provence, France (Richards, 1994); the Lower Permian Hueco Formation in south-central New Mexico (Mack and others, 2003); the Lower Cretaceous Woburn Sands in southern England (Yoshida and others, 2004); and the Eocene Central Basin in Spitsbergen (Plint-Björklund, 2005).

The base of the upper Cleveland TST is marked by the TS_30 marker, defined in many cores as the base of an upward-fining section that grades upward into either burrowed or carbonaceous siltstones. In electric logs, the base of the TST is inferred at inflection points in the gamma-ray (GR) curve at the base of upward-fining successions below the FS_30 flooding surface (Fig. 2). In the Shell #1-678 Wheat core, also presented in this workshop, no lowstand facies are preserved, and the transgressive systems tract is represented by a thin (commonly <3-ft [0.9-m]) section of fine-grained sandstone with ripple scours above an upward-coarsening section of highstand proximal-delta-front deposits (Fig. 3). The upper Cleveland TST is well developed in the Jones Energy #6 Tyson well, where it is an upward-fining 30-ft (9-m) section composed of a basal interval of sandy tidal-flat facies (Tstf) that grades upward into muddy tidal-flat (Tmtf) and marsh (Tm) facies.

Sandy and muddy tidal-flat facies (Tstf and Tmtf, respectively) in the upper Cleveland TST are herein described together because they have a similar genetic depositional origin, containing similar bedforms and reflecting similar depositional processes. However, they

are distinguished from one another by average grain size and general stratigraphic position, with the Tstf facies occurring lower in the TST succession (Fig. 2). The Tstf facies typically is composed of fine- to very fine grained sandstone with abundant ripples with mud drapes and internal scour surfaces. These ripples are dominantly asymmetric, although symmetrical, wavy ripples are also present. In contrast, the Tmtf facies is composed primarily of muddy siltstone and very fine grained sandstone. Stratification in the Tmtf facies is typified by abundant, small-scale, mud-draped ripples. Minor accessory features in the Tmtf facies include soft-sediment deformation, load structures at the bases of ripples, and carbonaceous fragments.

Upper Cleveland TST marsh facies are inferred from thin (commonly <1-cm) coal beds and carbonaceous shales. The marsh facies in the Jones Energy #6 Tyson well occurs at ~7,691 ft (~2,344.8 m) and overlies an 18-ft (5.5-m), upward-fining interval of Tmtf facies. In the Jones Energy #6 Tyson well it is composed of a 0.3-inch (0.7-cm), bright coal streak bounded below by gray mudstone with abundant plant fragments. It is overlain by muddy siltstone. In other cores from previous studies, marsh deposits are inferred from similarly thin (<4-inch [<10-cm]) coal seams that directly overlie upward-fining successions of crossbedded and ripple-laminated sandstone interpreted to be fluvial channel in origin (Hentz, 1994).

The upward-fining grain size in the tidal-flat and marsh section in the upper Cleveland TST sequence is consistent with overall grain-size profiles in prograding tidal-flat models in the Wash in the United Kingdom (Evans, 1965) and the southwestern coastline in the Netherlands (Van Stratten, 1954), as summarized in Klein (1971). In the prograding tidal-flat model from these modern examples, the upward-fining succession typically consists of a basal section of sandy, bedload-dominant, subtidal, and low-tidal-flat deposits that grade upward into finer-grained, middle- and upper-tidal-flat and marsh silt and mud deposits in which suspended-load sedimentation dominates over bedload sedimentation. Basinward superposition of muddy, upper-tidal-flat, and marsh deposits over lower-tidal-flat and subtidal deposits results in a net upward-fining grain-size profile. However, the upward-fining grain-size profile in transgressive tidal-flat and marsh deposits in the Jones Energy

#6 Tyson core is more a function of retrogradation. Muddy siltstones above the thin coaly zone in this core represent marine drowning (give-up surface) and destruction of marsh-forming environments. The give-up transgressive surface separates the lower peat deposit from the marine sediments above, dividing the two cycles into an upward-deepening cycle below and an upward-shallowing cycle above (Diessel et al., 2000).

#### **Cleveland Formation: Shell #1-678 Wheat Core**

#### **Highstand Systems Tract**

The Shell #1-678 Wheat core (Fig. 3) contains prominent, well-preserved highstand successions, in contrast to the Jones Energy #6 Tyson core. The lower highstand succession in the Shell #1-678 Wheat well from 7,331 to 7,404 ft (2,234 to 2,257 m) consists of a basal section of muddy inner-inner-shelf deposits (Hmis facies) that grades upward into distal-delta-front and proximal-delta-front sandstones (Hddf and Hpdf facies, respectively). This highstand succession is capped by a thin transgressive sandstone at 7,330 ft (2,234 m), in turn overlain by another upward-coarsening highstand succession.

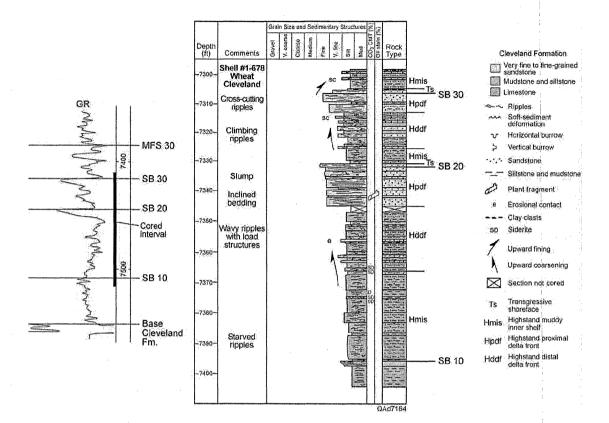


Figure 3. Description and log response of the Cleveland Formation in the Shell #1 Wheat core, located in Fig. 1.

Muddy inner-shelf facies (Hmis) composing the basal part of the highstand systems tract in the Cleveland Formation is overlain by distal-delta-front facies (Hddf). It consists of up to 30 ft (9 m) of sparsely burrowed, silty mudstone, with starved ripples and lenticular beds composed of fine-grained sandstone, as well as discontinuous siltstone laminae. These starved ripples and lenticular beds are typically <1 cm thick. The Hmis facies commonly exhibits no overall grain size, varying from fine- to coarse-grained siltstone, although it locally contains thin (<2-inch [<5.1-cm]) beds of calcareous, fine-grained sandstone.

The predominance of starved and isolated ripples in a muddy matrix in the Hmis facies indicates a low-energy shelf setting typified by suspension sedimentation and periodic sediment transport reflected by starved ripples, consistent with a sparse sand supply periodically transported by weak marine currents on a muddy substrate. The overall bedding is of the form described by Reineck and Wunderlich (1968) as lenticular bedding

with single lenses, in which >75% of the ripples (lenses) are discontinuous and encased in mudstone, stratification commonly associated with tidally influenced shelf environments. Weak wave reworking of these shelf deposits is recorded by minor, wavy, and symmetrical ripples, although these are small scale, with ripple heights commonly <0.5 inch (<1.3 cm).

Cleveland highstand delta-front deposits are divided into distal- (Hddf) and proximal-deltafront (Hpdf) facies. Net-sandstone maps of these highstand deltaic facies indicate that they are strongly dip-elongate, subparallel, and linear (see slide set "Sequence Stratigraphy and Depositional Summary of the Marmaton and Cleveland Formations, Anadarko Basin," this volume), consistent with net-sand patterns in modern subaqueous tidal-bar deposits such as those in the Gulf of Korea (Off, 1963), the Gulf of Papua offshore from the mouth of the Fly River (Fisher and others, 1969), and the Gulf of California (Meckel, 1975). A complete delta-front succession with both distal- and proximal-delta-front deposits is present in the Shell #1-678 Wheat well from 7,331 to 7,366 ft (2,234 to 2,245 m). The distal-delta-front section, inferred from 7.345 to 7.366 ft (2.239 to 2.245 m), is composed of sparsely burrowed siltstone and thin (<3-inch [<7.6-cm]) beds of very fine grained sandstone. Sedimentary structures in these distal-delta-front deposits are dominated by asymmetric and wavy, starved ripples. Soft-sediment deformation is common, occurring as slumps and load structures. The proximal-delta-front facies (Hpdf) in the Shell #1-678 Wheat core abruptly overlies the distal-delta-front facies and is composed of upper-fine-grained sandstone. Stratification in the proximal-delta-front facies in the Shell #1-678 Wheat well consists of slightly inclined laminae and crossbeds. Soft-sediment deformation is also present, occurring as slumps with overturned strata. Slumps and soft-sediment deformation, although common in delta-front settings in fluvial-dominated deltas (Coleman and others, 1974; Prior and Coleman, 1978; Elliott, 1989), can also occur in tide-dominated facies, where rapid deposition of material from flood-tidal currents on an unstable substrate can result in rotation and failure of semicoherent material (Rahman and others, 2008; Carmona and others, 2009). The top of the Hpdf facies is capped by a thin, upper, fine-grained sandstone at 7,331 ft (2,234 m), interpreted to represent transgressive deposits, which is in turn overlain by muddy inner-shelf deposits in another highstand cycle. This second highstand succession, which ranges from 7,307 to 7,331 ft (2,227 to 2,234 m), is an upward-coarsening interval that ranges from lenticular, starved ripples in a muddy matrix at the base, grading upward into large-scale, climbing, wavy ripples with significant mud drapes. The section is capped by another thin (<6-inch [<15-cm]) zone of transgressive deposits at 7,307 ft (2,227 m), identified as fine- to very fine grained sandstone with abundant internal scours. This transgressive sandstone is abruptly overlain by burrowed, silty mudstone with thin (millimeter-scale) laminae and starved ripples composed of very fine grained sandstone. This fine-grained section is the lower part of another highstand succession at the top of the cored interval.

#### Marmaton Group: Sun #1 Blau Core

A small, 22-ft (6.7-m) cored section from the Sun #1 Blau core, located in Fig. 1, is included to show sedimentary structures and facies within a highstand succession in the Marmaton Group for the sake of comparison with similar cored intervals in the Cleveland Formation. This cored interval is from an upward-coarsening Marmaton parasequence that grades upward from horizontally stratified, very fine grained sandstone at the base to finegrained sandstones with large, wavy ripples with mudstone drapes that are in turn overlain by fine-grained sandstones with internal scour surfaces and flattened clay clasts (Fig. 4). The main part of the cored section from 7,548 to 7,562 ft (2,301 to 2,305 m) is composed of draped, wavy ripples that vary in type from starved ripples/flasers at 7,560 ft (2,304 m) to shale-draped, climbing ripples and large, climbing ripples from 7,552 to 7,560 ft (2,302 to 2,304 m). Many of the shale drapes occur as couplets, especially at 7,554 to 7,555 ft (2,302 to 2,303 m). The upper one-third of the section is dominated by fine-grained sandstone with clay clasts and ripple scours, suggesting relatively high energy deposition and erosional processes consistent with wave or storm scouring. Although the spontaneous potential (SP) curve in the cored section is nonresponsive, the resistivity curve suggests an upward-coarsening grain-size profile. The top of the cored section is near the sequence boundary M_SB_20, inferred in the core at ~7,545 ft (~2,300 m) along the base of the uppermost sandstone bed with internal scour surfaces and clay clasts. The overall upwardcoarsening grain-size profile in the cored section is interpreted to record a progradation and offlap of a tidally influenced delta-front succession capped by ravinement scouring during

a regional transgression subsequent to a relative-sea-level fall marked by the M_SB_20 sequence boundary.

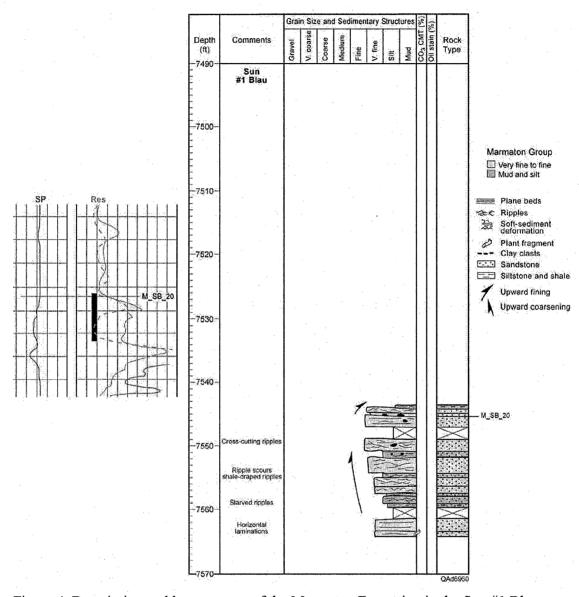


Figure 4. Description and log response of the Marmaton Formation in the Sun #1 Blau core.

#### References

- Allen, G. P., 1991, Sedimentary processes and facies in the Gironde estuary: a Recent model of macrotidal estuarine systems, *in* G. D. Smith, G. E. Reinson, B. A. Zaitlin, and R. A. Rahmani, eds., Clastic Tidal Sedimentology: Canadian Society of Petroleum Geologists Memoir 16, p. 29–40.
- Allen, G. P., and H. W. Posamentier, 1993, Sequence stratigraphy and facies model of an incised valley fill: the Gironde Estuary, France: Journal of Sedimentary Petrology, v. 63, no. 3, p. 378–391.
- Ambrose, W. A., E. R. Ferrer, S. P. Dutton, D. L. Wang, A. Padron, W. Carrasquel, J. S. Yeh, and N. Tyler, 1995, Production optimization of tide-dominated deltaic reservoirs of the Lower Misoa Formation (lower Eocene), LL-652 Area, Lagunillas field, Lake Maracaibo, Venezuela: The University of Texas at Austin, Bureau of Economic Geology Report of Investigations No. 226, 46 p.
- Bridges, P. H, and M. R. Leeder, 1976, Sedimentary model for intertidal mudflat channels with examples from the Solway Firth, Scotland: Sedimentology, v. 23, p. 533–552.
- Carmona, N. B., L. A. Buatois, J. J. Ponce, and M. G. Mángano, 2009, Ichnology and sedimentology of a tide-influenced delta, Lower Miocene Chenque Formation, Patagonia, Argentina: Trace-fossil distribution and response to environmental stresses: Palaeogeography, Palaeoclimatology, Palaeoecology, v. 273, p. 75–86.

- Coleman, J. M., J. N. Suhayda, T. Whelan, and L. D. Wright, 1974, Mass movement of Mississippi River sediments: Gulf Coast Association of Geological Societies Transactions, v. 24, p. 49-68.
- Dalrymple, R. W., 1992, Tidal depositional systems, *in* R. G. Walker and N. P. James, eds.,Facies Models: Response to Sea Level Changes: Geological Association of Canada,p. 195–218.
- Dalrymple, R. W., R. J. Knight, B. A. Zaitlin, and G. V. Middleton, 1990, Dynamics and facies model of a macrotidal sand-bar complex, Cobequid Bay-Salmon River estuary (Bay of Fundy): Sedimentology, v. 37, p. 577–612.
- Dalrymple, R. W., Y. Makino, and B. A. Zaitlin, 1991, Temporal and spatial patterns of rhythmite deposition on mudflats in the macrotidal Cobequid Bay-Salmon River estuary, Bay of Fundy, Canada, *in* D. G. Smith, G. E. Reinson, B. A. Zaitlin, and R. A. Rahmani, eds., Clastic Tidal Sedimentology: Canadian Society of Petroleum Geologists Memoir 16, p. 137–160.
- Dalrymple, R. W., B. A. Zaitlin, and R. Boyd, 1992, Estuarine facies models: conceptual basis and stratigraphic implications: Journal of Sedimentary Petrology, v. 62, no. 6, p. 1130–1146.

Davies, J. L., 1964, A morphogenic approach to world shorelines: Zeitschrift Für Geomorphologie, Band 8, p. 27–42.

- de Mowbray, Tessa, and Visser, M. J., 1984, Reactivation surfaces in subtidal channel deposits, Oosterschelde, southwest Netherlands: Journal of Sedimentary Petrology, v. 54, no. 3, p. 811–824.
- Devine, P. E., 1991, Transgressive origin of channelized estuarine deposits in the Point Lookout Sandstone, northwestern New Mexico: A model for Upper Cretaceous, cyclic regressive parasequences of the U.S. Western Interior: AAPG Bulletin, v. 75, no. 6, p. 1039–1063.
- Diessel, C., R. Boyd, J. Wadsworth, D. Leckie, and G. Chalmers, 2000, On balanced and unbalanced accommodation/peat accumulation ratios in the Cretaceous coals from Gates Formation, Western Canada, and their sequence-stratigraphic significance: International Journal of Coal Geology, v. 43, p.143–186.
- Elliott, T., 1989, Deltaic systems and their contribution to an understanding of basin-fill successions, *in* M. K. G. Whateley and K. T. Pickering, eds., Deltas: Sites and Traps for Fossil Fuels: Geological Society of London Special Publication No. 41, p. 3–10.
- Evans, G., 1965, Intertidal flat sediments and their environments of deposition in the wash: Geological Society of London Quarterly Journal, v.121, p. 209–245.
- Fisher, W. L., L. F. Brown, Jr., A. J. Scott, and J. H. McGowen, 1969, Delta systems in the exploration for oil and gas: The University of Texas at Austin, Bureau of Economic Geology, Research Colloquium, variously paginated.

- Greer, S. A., 1975, Sandbody geometry and sedimentary facies at the estuary-marine transition zone, Ossabow Sound, Georgia: A stratigraphic model: Senckenbergiana Maritima, Band 7, p. 105–136.
- Hayes, M. O., 1976, Morphology of sand accumulation in estuaries: an introduction to the symposium, *in* L. E. Cronin, ed., Estuarine Research, v. II: Geology and Engineering, Academic Press, New York, p. 3–22.
- Hentz, T. F., 1994, Depositional, structural, and sequence framework of the gas-bearing Cleveland Formation (Upper Pennsylvanian), western Anadarko Basin, Texas Panhandle: The University of Texas at Austin, Bureau of Economic Geology Report of Investigations No. 213, 73 p.
- Hentz, T. F., W. A. Ambrose, and D. L. Carr, 2009, Sequence Stratigraphic and Depositional Settings of Highstand Deltaic and Lowstand Valley-Fill Deposits of the Middle and Upper Pennsylvanian Cleveland and Marmaton Tight-Gas Sandstones, Northwest Anadarko Basin (abs.): AAPG Search and Discovery Article #90090.
- Houbolt, J. J. H. C., 1968, Recent sediments in the southern bight of the North Sea: Geologie en Mijnbouw, v. 47, p. 245–273.
- Houthuys, R., and F. Gullentops, 1988, The Vlierzele Sands (Eocene, Belgium): a tidal ridge system, in P. L. de Boer, A. van Gelder, and S. D. Nio, eds., Tide-Influenced Sedimentary Environments and Facies: D. Reidel Publishing Company, Boston, p. 139–152.

- Klein, G. de V., 1970, Depositional and dispersal dynamics of intertidal sand bars: Journal of Sedimentary Petrology, v. 40, p. 1095–1127.
- Klein, G. de V., 1971, A sedimentary model for determining paleotidal range: Geological Society of America Bulletin, v. 82, p. 2585–2592.
- Kreisa, R. D., and R. J. Miolola, 1986, Sigmoidal tidal bundles and other tide-generated sedimentary structures of the Curtis Formation, Utah: Geological Society of America Bulletin, v. 97, p. 381–387.
- Lambiase, J. J., 1980, Hydraulic control of grain-size distribution in a macrotidal estuary: Sedimentology, v. 27, p. 433–446.
- Leckie, D. A., and C. Singh, 1991, Estuarine deposits of the Albian Paddy Member (Peace River Formation) and Lowermost Shaftesbury Formation, Alberta, Canada: Journal of Sedimentary Petrology, v. 61, no. 5, p. 825–849.
- Mack, G. H., M. Leeder, M. Perez-Arlucea, and B. D. J. Bailey, 2003, Sedimentology, paleontology, and sequence stratigraphy of early Permian estuarine deposits, south-central New Mexico, USA: Palaios, v. 18, no. 4-5, p. 403-420.
- Maguregui, J. A., and N. Tyler, 1991, Evolution of middle Eocene tide-dominated deltaic sandstones, Lagunillas field, Maracaibo Basin, western Venezuela, *in* A. D. Miall, and N. Tyler, eds., The Three-Dimensional Facies Architecture of Terrigenous Clastic

Sediments and Its Implications for Hydrocarbon Discovery and Recovery: SEPM (Society for Sedimentary Geology), v. 3, Concepts in Sedimentology and Paleontology, p. 233–244.

- Meckel, L. D., 1975, Holocene sand bodies in the Colorado Delta area, northern Gulf of California, *in* M. L. Broussard, ed., Deltas, models for exploration: Houston Geological Society, p. 239–265.
- Mutti, E., J. Rossell, G. P. Allen, F. Fonnesu, and M. Sgavetti, 1985, The Eocene Baronia tide dominated delta-shelf system in the Ager Basin, *in* M. D. Miall, and J. Rossell, eds., Excursion Guidebook: 6th European Regional Meeting, International Association of Sedimentologists, Lleida, Spain, p. 579–600.
- Off, T., 1963, Rhythmic linear sand bodies caused by tidal currents: AAPG Bulletin, v. 47, no. 2, p. 324–341.
- Plint-Björklund, P., 2005, Stacked fluvial and tide-dominated estuarine deposits in highfrequency (fourth-order) sequences of the Eocene Central Basin, Spitsbergen: Sedimentology, v. 52, p. 391–428.
- Prior, D. B., and J. M. Coleman, 1978, Submarine landslides on the Mississippi river deltafront slope: Geoscience and Man, v. 19, p. 41–53.
- Rahman, M. J. J., P. Faupl, and M. M. Alam, 2008, Depositional facies of the subsurface Neogene Surma Group in the Sylhet Trough of the Bengal Basin, Bangladesh: Record of tidal sedimentation: International Journal of Earth Sciences, DOI 10.1007/s00531-008-0347-7.

- Reineck, H. E., 1967, Layered sediments of tidal flats, beaches and shelf bottoms of the North Sea, *in* G. D. Lauff, ed., Estuaries: American Association for the Advancement of Science, p. 191–206.
- Reineck, H. E., and F. Wunderlich, 1968, Classification and origin of flaser and lenticular bedding: Sedimentology, v. 11, p. 99–104.
- Reineck, H. E., and Singh, I. B., 1973, Depositional sedimentary environments: with reference to terrigenous clastics: Springer-Verlag, Berlin, 439 p.
- Richards, M. T., 1994, Transgression of an estuarine channel and tidal flat complex: The lower Triassic of Barles, Alpes de Haute Provence, France: Sedimentology, v. 41, p. 55–82.
- Swift, D. J. P., 1975, Tidal sand ridges and shoal-retreat massifs: Marine Geology, v. 18, p. 105–134.
- Van Stratten, L. M. J. U., 1954, Composition and structure of recent marine sediments in the Netherlands: Leidse Geologische Mededelingen, v. 19, p. 1–110.
- Wright, L. D., J. M. Coleman, and B. G. Thom, 1973, Processes of channel development in a high tide environment: Cambridge Gulf, Ord River Delta, western Australia: Journal of Geology, v. 81, p. 15–41.
- Wright, L. D., J. M. Coleman, and B. G. Thom, 1976, Sediment transport and deposition in a macrotidal river channel: Ord River, western Australia, *in* L. E. Cronin, ed.:

Estuarine Research, v. II: Geology and Engineering, Academic Press, New York, p. 309–322.

Yoshida, S., H. D. Johnson, K. Pye, and R. J. Dixon, 2004, Transgressive changes from tidal estuarine to marine embayment depositional systems: The Lower Cretaceous Woburn Sands of southern England and comparison with Holocene analogs: AAPG Bulletin, v. 88, no. 10, p. 1433–1460.

PITC WESTAPS bandouts for ARPG

# electronic?

Plectro	MC .		A			· · · · · ·	
N N	Southwest Southwest	2002 2002	9/17/2002 9/26/2002	9/17/2002 9/26/2002	Wellbore Management New Resources from Old Fields: Revitalizing Gas Exploration and Production in the GOM Shelf	Tyler Houston	TX TX
N	Southwest	2003	2/27/03	2/27/2003	Electronic Resources for NM Producers	Midland	ТХ
N	Southwest	2003	5/19/03	5/19/2003	H2S Safety Seminar (NM Oil Conservation Division, WERC)	Midland	TX
N N N	Southwest Texas Texas	2005 1996 1997	5/12/05 4/8/1996 10/24/1997	5/12/2005 4/8/1996 10/24/1997	Horizontal Drilling (Texas) Frio Fluvial - Deltaic Sandstone Play Effect of Ellenburger Karsting & Collapse on Reservoir Compartmentalization	Midland San Antonio San Antonio	TX TX TX
N N N	Texas Texas Texas Texas	1997 1998 1999 1999	11/13/1997 4/23/1998 9/16/1999 10/18/1999	11/13/1997 4/23/1998 9/16/1999 10/18/99 to 10/21/99	Advanced Logging Techniques 3-D Seismic (Waha/Lockridge) Mudlogging 3D Geophysics for Geologists (E. Tex. Geol. Soc.)	Midland Midland Midland Tyler	TX TX TX TX
N.	Texas Texas	2000 2000	2/17/2000 3/9/2000	2/17/2000 3/9/2000	Vapor Recovery Workshop Coalbed Methane Workshop (Houston Geol. Society)	Midland Houston	TX TX
N,	Texas	2003	4/10/03	04/10-11/03	Structure and Stratigraphy of South Texas and Northeast Mexico - Applications to Exploration (STGS, Gulf Coast SEPM)	Austin	TX
N 1	Texas T _e xas	2003 2003	4/15/03 5/29/03	4/15/2003 5/29/2003	Well Cuttings New Methods for Locating and Recovery Remaining Hydrocarbons in the Permian Basin (UT Bureau of Economics Geology, Center for Energy and Economic Diversification)	Houston Midland	TX TX
N	Texas	2004	1/15/04	1/15/2004	Integrated Synthesis of Permian Basin Depositional Systems	Midland	TX
N	Texas	2004	5/18/04	5/18/2004	Field Data Entry (Panhandle Producers and Royalty Association)	Amarillo	ТХ
N	Texas	2004	10/28/04	10/28/2004	Tech Session @ Permian Basin Petroleum Association annual mtg	Midland	ТХ
M	Texas	2004	11/4/04	11/4/2004	Produced Water Management (Texas Alliance)	Wichita Falls	тх
e se x	Texas	2005	3/1/05	3/1/2005	Data Gathering Techniques & Interfaced Production Accounting Software	Midland	ТХ
	Texas Texas Texas	2005 2005 2005	3/2/05 6/29/05 9/20/05	3/2/2005 6/29/2005 9/20/2005	Controlling Sand Production Reservoir Engineering Stranded Gas, Options for Realizing Value (Texas Alliance of Energy Producers)	Houston Houston Dallas	TX TX TX
	Texas	2005	10/20/05	10/20/2005	Tech Session @ Permian Basin Petroleum Association Annual Meeting	Midland	ТΧ
. (	Téxas	2006	8/1/06	8/1/2006	Forecasting Waterfloods Using Reservoir Grail (Grail Quest, Midland College)	Midland	ТХ
					- <i>,</i>		

	(Texas	Sarahe					
yer	Texas Dor	2006	11/8/06	11/8/2006	Barnett Shale Gas Play of the Fort Worth Basin (Bureau of Economic Geology)	Midland	ТХ
yes	Texas	2007	2/23/07	2/23/2007	Hydraulic Fracturing Technology and Case Studies, Tight Gas Sands and Shales (Texas Alliance, South Texas Geological Society)	San Antonio	ТХ
Bob	Texas	2007	10/17/07	10/17/2007	Production Data Gathering and Remote Surveillance	Midland	тх
yes	Texas	2007	10/30/07	10/30/2007	Pilot Study of the East Texas Field; Geology, Engineering and Potential Future Exploitation	Kilgore	ТХ
VI	Texas/Central Gulf	1998	12/1/1998	12/1/1998	Fractured Reservoirs (w/DOE & PTTC Central Gulf)	Tyler	тх

### PTTC Texas/ SE New Mexico Region Workshops, 2000-2008

How to Start/Fix/Manage a Small Waterflood and Successful Practices, Midland, Texas, May 2008

Technologies for Developing Naturally Fractured Reservoirs, Houston, Texas, February 2008

Designing and Forecasting Waterfloods Using "ReservoirGrail" "The Best Place to Find Oil is in Oil Fields!", Farmers Branch, Texas, November 2007

Pilot Study of the East Texas Field: Geology, Engineering, and Potential Exploitation, Kilgore, Texas, October 2007

Production Data Gathering and Remote Surveillance, Midland, Texas, October 2007

Hydraulic Fracturing Technology and Case Studies, Tight Gas Sands and Shales, San Antonio, Texas, February 2007

Barnett Shale-Gas Play of the Fort Worth Basin, Austin, Texas, November 2006

Introduction to Oilfield Explosives Safety Seminar, PTTC Texas Region, Houston, Texas, August 2006

Hydraulic Fracturing Technology and Case Studies—Tight Gas Sands and Shales, PTTC Texas Region, Irving, Texas, July 2006

Designing and Forecasting Waterfloods Using "Reservoir Grail"—"The Best Place to Find Oil is in Oil Fields!" PTTC Texas Region, Midland, Texas, August 2006

Hydraulic Fracturing Technology and Case Studies—Tight Gas Sands and Shales, PTTC Texas Region, Tyler, Texas, July 2006

Producers Technology Transfer Workshop—Natural Gas STAR Program, PTTC Texas Region, Midland, Texas, June 2006

Producers Technology Transfer Workshop—Natural Gas STAR Program, PTTC Texas Region, Fort Worth, Texas, June 2006

Petroleum Geoscience: Basics/ of Petroleum Generation, Migration, Trapping and the Oil Business, SIPES Dallas Chapter and PTTC Texas Region, Farmer's Branch, Texas, May 2006

West Texas Barnett Shale Symposium, Midland College and PTTC Texas Region, Midland, Texas, April 2006

Hydraulic Fracturing Technology and Case Studies in Permian Basin Tight Gas Sands and Shales, Texas Alliance of Energy Producers and PTTC Texas Region, Midland, Texas, March 2006

The Tight Gas Sands of the Cotton Valley Formation of East Texas: a Core Workshop, Bo Henks and PTTC Texas Region, Austin, Texas, March 2006

Horizontal Well Technologies and Coalbed Methane Applications, PTTC Central Gulf and Texas Regions, Houston, Texas, February 2006

The Tight Gas Sands of the Cotton Valley Formation of East Texas: A Core Workshop: presented to PTTC Texas Region, Austin, Texas, January 2006.

The Tight Gas Sands of the Cotton Valley Formation of East Texas: A Core Workshop: presented to PTTC Texas Region, Austin, Texas, December 2005.

Reserve Growth Potential from CO₂ Enhanced Oil Recovery along the Gulf Coast: presented to PTTC Texas Region, Houston, Texas, December 2005.

Fundamentals of Seismic Interpretation: presented to PTTC Texas Region, Farmers Branch, Texas, November 2005.

Symposium on Improved Profits through Best Managed Practice: presented to PTTC Texas Region, Corpus Christi, Texas, November 2005.

Barnett Shale Symposium: presented to PTTC Texas Region, Midland, Texas, November 2005.

New Technology Enabling New Plays: presented to PTTC Texas Region, Farmers Branch, Texas, October 2005.

Geological, Petrophysical, and Engineering Aspects of Avoiding Reserves Writedowns: presented to PTTC Texas Region, Farmers Branch, Texas, September 2005.

Stranded Gas, Options for Realizing Value: presented to PTTC Texas Region, Wichita Falls, Texas, September 2005.

Introduction to Mining the Internet: Using Free GIS Data and Low Cost Software for the Oil & Gas Professional: presented to PTTC Texas Region, Midland, Texas, September 2005.

Introduction to Mining the Internet: Using Free GIS Data and Low Cost Software for the Oil & Gas Professional: presented to PTTC Texas Region, Farmington, New Mexico, September 2005.

Barnett Shale Symposium III: presented to PTTC Texas Region, Farmers Branch, Texas, June 2005.

Turning Emissions into Significant Profits: presented to PTTC Texas Region, Corpus Christi, Texas, June 2005.

Material Balance, Modeling, and Simulation: Reservoir Engineering Tools Past, Present, and Future: presented to PTTC Texas Region, Houston, Texas, June 2005.

Drillinginfo.com seminar: presented to PTTC Texas Region, Farmers Branch, Texas, May 2005.

Rocks, Pores, and Capillary Pressure: The Hole Story on How to Understand Reservoirs and Seals by Thinking like Oil and Gas: presented to PTTC Texas Region, Farmers Branch, Texas, March 2005.

Production Data: Collecting It and Using It!: presented to PTTC Texas Region, Midland, Texas, March 2005.

Production Data: Collecting It and Using It!: presented to PTTC Texas Region, Farmington, New Mexico, March 2005.

Sand Control Workshop: presented to PTTC Texas Region, Houston, Texas, March 2005.

Stratigraphic Synthesis of Paleozoic Oil-Bearing Depositional Systems: Data and Models for Recovering Existing and Undiscovered Oil Resources from the Permian Basin: presented to PTTC Texas Region, Houston, Texas, December 2004.

Lunch and Learn: Technology in Well Optimization Utilizing New Technologies: presented to PTTC Texas Region, Houston, Texas, December 2004.

Converting Stranded Gas and Lowering Power/Operating Costs Workshop: presented to PTTC Texas Region, Midland, Texas, December 2004.

Deep Gas Well Simulation Workshop: presented to PTTC Texas Region, Houston, Texas, December 2004.

Lunch and Learn: Technology in Well Optimization Utilizing New Technologies in Field Automation and Remote Monitoring: presented to PTTC Texas Region, Midland, Texas, December 2004.

Essentials of Subsurface Mapping: presented to PTTC Texas Region, San Antonio, Texas, November 2004.

Stratigraphic Synthesis of Paleozoic Oil-Bearing Depositional Systems: Data and Models for Recovering Existing and Undiscovered Oil Resources from the Permian Basin: presented to PTTC Texas Region, Midland, Texas, November 2004.

Unconventional Reservoirs Symposium: presented to PTTC Texas Region, Farmers Branch, Texas, November 2004.

Produced Water and Associated Issues: presented to PTTC Texas Region, Wichita Falls, Texas, November 2004.

Polymer and Polymer-Gel Water Shutoff Treatments: What It Takes To Be Successful and Illustrative Field Applications: PTTC Texas Region, Houston, Texas, August 2004.

Reservoir Fluids: from the Matrix to the Market—What You Don't Know CAN Hurt You: PTTC Texas Region, Houston, Texas, July 2004.

Operators Working in Offshore Texas State Waters Royalty Leases: PTTC Texas Region, Houston, Texas, June 2004.

Soil Remediation: PTTC Texas Region, Hobbs, New Mexico, June 2004.

Barnett Shale and Other Fort Worth Basin Plays: PTTC Texas Region, Farmers Branch, Texas, June 2004.

Low-Cost Rod Pump Control: PTTC Texas Region, Farmers Branch, Texas, March 2004.

Low-Cost Rod Pump Control: PTTC Texas Region, Midland, Texas, March 2004.

Soil Remediation: PTTC Texas Region, Tyler, Texas, January 2004.

Integrated Synthesis of Permian Basin Depositional History and Stratigraphic Architecture: Data and Models for Recovering Existing and Undiscovered Oil Resources: PTTC Texas Region, Midland, Texas, January 2004.

Low-Cost Rod Pump Control: PTTC Texas Region, Houston, Texas, October 2003.

Understanding Paraffin and Asphaltene Problems in Oil and Gas Wells: PTTC Texas Region, San Angelo, Texas, September 2003.

Visualization Seminar: presented at Bureau of Economic Geology, Austin, Texas, September 2003.

Inexpensive, Rapid Cross-Section Generation Utilizing MJ Systems' Raster Log Images and DigiRule's CrossLog Suite[™] Software: PTTC Texas Region, Austin, Texas, August 2003.

Produced Water and Associated Issues: PTTC Texas Region, Farmers Branch, Texas, August 2003.

Produced Water and Associated Issues: PTTC Texas Region, Midland, Texas, August 2003.

Reservoir Fluids: PTTC Texas Region, Houston, Texas, June 2003.

New Methods for Locating and Recovering Remaining Hydrocarbons in the Permian Basin: PTTC Texas Region, Midland, Texas, May 2003.

Hydraulic Fracturing: PTTC Texas Region, Shreveport, Louisiana, April 2003.

Well Cuttings: PTTC Texas Region, Houston Research Center, Houston, Texas, April 2003.

SONRIS Database Web Interface: PTTC, Houston, Texas, November 2002.

Inexpensive, Rapid Cross-section Generation Utilizing MJ Systems' Raster Log Images and DigiRule's CrossLog Suite™ Software: PTTC Texas Region, Austin, Texas, October 2002.

New Resources from Old Fields: Revitalizing Gas Exploration and Production in the Gulf of Mexico Shelf Province: PTTC Texas Region, Austin, Texas, September 2002.

Wellbore Management: PTTC Texas Region, Tyler, Texas, September 2002.

Inexpensive, Rapid Cross-section Generation Utilizing MJ Systems' Raster Log Images and DigiRule's CrossLog Suite™ Software: PTTC Texas Region, Austin, Texas, September 2002.

Optimizing Horizontal Well Technology: PTTC Texas Region, Austin, Texas, May 2001.

Predicting Reservoir Quality Using Diagenetic Models: PTTC Texas Region, Austin, Texas, March 2001.

Reservoir Characterization Technologies for the Next Millennium: Virtual Reality, Multicomponent Seismic, Fracture Modeling, and Borehole Imaging: PTTC Texas Region, Austin, Texas, February 2000.