



Effect of Geometry and Number of Seismic Stations on Micro-Earthquake (MEQ) Hypocenters in Geothermal Fields

Widya Utama^{1*}, Sherly Ardhya Garini², Merry C. Hutapea¹, Dhea Pratama Novian Putra³, Dwa Desa Warnana¹, Wien Lestari¹

¹ Department of Geophysics Engineering, Institut Teknologi Sepuluh Nopember, Surabaya 60111, Indonesia.

² Department of Informatics, Institut Teknologi Sepuluh Nopember, Surabaya 60111, Indonesia.

³ Department of Geomatics Engineering, Institut Teknologi Sepuluh Nopember, Surabaya 60111, Indonesia.

Received: April 28, 2023

Revised: September 15, 2023

Accepted: October 25, 2023

Published: October 31, 2023

Corresponding Author:

Widya Utama

widya@geofisika.its.ac.id

DOI: [10.29303/jppipa.v9i10.3742](https://doi.org/10.29303/jppipa.v9i10.3742)

© 2023 The Authors. This open access article is distributed under a (CC-BY License)



Abstract: Micro-earthquake (MEQ) distribution describes subsurface conditions that can contribute to monitoring the dynamics of geothermal reservoirs. Thus, the distribution of MEQ hypocenter locations with high accuracy becomes extremely important. Experiments were conducted with 3 variations of geometry and number of seismic stations, while Geiger and Coupled Velocity-Hypocenter methods were used to determine the location of MEQ. Experimental results show that in determining the location of the MEQ, the geometry and number of seismic stations played an important role. Increasing the number of stations with relatively long distances can result in less accurate locations of MEQ, error and bias in determining the location of MEQ will be greater when the azimuth gap value is greater. This is shown by the distribution of MEQ that are more spread out in variations 4A and 4B (4 seismic stations) compared to the distribution of MEQ hypocenters using data from 8 seismic stations. The azimuth gap variations of stations 4A and 4B are 283° and 267°, and 8 stations have a value of 222°. The large value of the azimuth gap is due to the distribution of stations only on one side so that there are horizontal angles that are not covered by seismic stations.

Keywords: Geiger; Geometry; Geothermal; Micro-earthquake (MEQ); Seismic Station.

Introduction

Geothermal is one of the natural resources formed from high-temperature materials from under the surface of the earth that is used by humans to provide for their daily needs (Aneke & Menkiti, 2016; Bertani, 2016; Cheng, 2022; Lund et al., 2011; Rahmanningtyas et al., 2020). Some geothermal energy uses as power generation to heating necessity (Jouhara et al., 2020; Moya et al., 2018). Geothermal energy is included in environmentally friendly energy (Anderson & Rezaie, 2019; Østergaard et al., 2020). The utilization of geothermal energy is included in the Sustainable Development Goals (SDGs) as stated in point 7 (related to clean and affordable energy), point 9 (related to industry, innovation, and infrastructure), point 11 (related to sustainable cities and communities), point 13

(related to climate action), and point 17 (related to partnerships for goals).

The existence of geothermal energy is generally found in the volcanic zone, which is a weak zone in geology (Cheng, 2022; Farhan et al., 2019; Geoffroy et al., 2022; Mahwa et al., 2022). In this zone, there are many faults formed due to geological activity (Geoffroy et al., 2022; Riziq Maulana et al., 2019; Zaini et al., 2022). The faults that are formed open a fluid migration path (generally groundwater) to approach the heat source and or migrate to the geothermal reservoir zone (Anderson & Rezaie, 2019; Chen & Huang, 2018; Cheng, 2022; Geoffroy et al., 2022). Fault planes will generally increase porosity and permeability due to fractures formed between one fault plane and another (Chen & Huang, 2018; Geoffroy et al., 2022; Riziq Maulana et al., 2019; Yang et al., 2021). The presence of faults that cause

How to Cite:

Utama, W., Garini, S. A., Hutapea, M. C., Putra, D. P. N., Warnana, D. D., & Lestari, W. (2023). Effect of Geometry and Number of Seismic Stations on Micro-Earthquake (MEQ) Hypocenters in Geothermal Fields. *Jurnal Penelitian Pendidikan IPA*, 9(10), 8114-8123. <https://doi.org/10.29303/jppipa.v9i10.3742>

this fracturing can be identified based on the release of seismic waves generated when the fault is formed (Barbosa et al., 2020; Farhan et al., 2019; Firdaus Al Hakim et al., 2019; Küperkoch et al., 2018; Riziq Maulana et al., 2019).

The Micro-Earthquake (MEQ) method can be used to identify the presence of faults, where the distribution of MEQ hypocenters with a relatively low magnitude scale, i.e. with a strength of less than 3 Richter Scale (SR) can indicate the presence of faults (Kato et al., 2021; Toledo et al., 2020; Utama et al., 2021; Wildan Perdana et al., 2020). MEQ is used to identify subsurface dynamics related to the presence of faults and weak zones in geothermal field reservoirs, MEQ hypocenter determination is carried out based on seismic data recorded by seismic stations which are then processed by the inversion method and then analyzed to describe the subsurface conditions of the geothermal field (Pennington et al., 2022; Utama, Ardhy, et al., 2022). Generally, after the identification of the hypocenter is carried out, analysis and validation of the suitability with geological data of the geothermal field area or research area are carried out (Halim et al., 2020; Intani et al., 2020; Sicking & Malin, 2019; Utama & Garini, 2022).

Based on the results of previous studies, the quality of reliable MEQ location with a high level of precision and accuracy depends on several factors, including the availability of good quality (low noise) data recorded by the cluster network, the measurement error of the observed arrival time (quality of seismic phase determination, number and distribution of seismic stations, subsurface seismic velocity structure in the cluster region), the quality of seismic phase determination, the number and distribution of seismic stations, and the accuracy of MEQ subsurface seismic velocity in the unknown cluster region (Karasözen and Karasözen 2020; Kianimehr et al. 2018; Midzi et al. 2020; Sevilla et al. 2020; Zhang et al., 2019; Utama, Garini, and Lansa 2022).

Research on the effect of the number of seismic stations on the estimation of the earthquake Centroid Moment Tensor (CMT) has been widely conducted, where the number of seismic stations affects the estimation of the earthquake CMT (Fahntalia & Madlazim, 2017). The influence of the number and distribution of seismic stations on the identification of MEQ hypocenter locations in geothermal fields is something that has been studied (Bidang, 2020; Cheng, 2022; Fahntalia & Madlazim, 2017; Pennington et al., 2022). There are indications that the number and geometry of seismic stations influence the distribution of MEQs, thus affecting the analysis results obtained. Therefore, some MEQ data processing software applies a minimum number of stations used for processing, in

order to obtain a more representative MEQ hypocenter model (Kissling, 1998; Kissling et al., 1995; Nishi, 2005). When the number of stations used does not reach the minimum limit, it will produce a relatively large error value, resulting in less accurate identification of the hypocenter (Bondár et al., 2004; Huang et al., 2018; Küperkoch et al., 2018; Siddiq et al., 2020).

This research applies the Geiger Method with Adaptive Damping (GAD) and Coupled Velocity Hypocenter (CVH) to identify MEQ hypocenters in geothermal fields, related to the influence of the number of seismic stations and their location distribution. Variations in the number of seismic stations used in this study are 4 and 8 seismic stations. The subsurface rock model in this study is considered a homogeneous model, so that the variation of rock physical properties as a medium for MEQ seismic wave propagation is considered isotropic. This research aims to support the 7th point of SDGs in realizing clean and affordable energy, through geothermal energy applications that require detailed fault conditions and locations, which can be identified from the studied MEQ hypocenter.

Method

The number of stations used in the analysis of the influence of the number of seismic stations and distribution on the identification of MEQ hypocenter locations are 4 and 8 seismic stations, namely ULI, ZPN, ZUI, BRB, BUT, BRK, ZTO, and BUA seismic stations. In analyzing the influence of hypocenter distribution, 3 types of variations are used, namely 4A, 4B and 8 to get an overview of the differences between the two approaches. The first variation 4A consists of 4 stations which are ULI, ZPN, ZTO, and BUT. The 4B variation consists of stations ZUI, BRK, BRB, and BUA, and the third variation consists of all stations namely ULI, ZPN, ZTO, BUT, ZUI, BRK, BRB, and BUA. There are several input data needed to identify the location of MEQ using the inversion method including the arrival time of Primary (P) and Secondary (S) waves, seismic station coordinates, magnitude, origin time, and 1D velocity model. The inversion methods used are Geiger method to identify the initial MEQ location and the Coupled Velocity-Hypocenter (CVH) method to relocate the MEQ and determine the 1D velocity model of the cluster region, the framework can be seen in Figure 1.

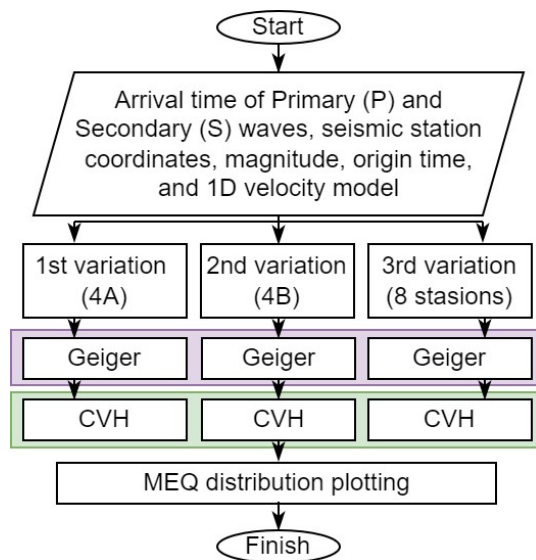


Figure 1. Research flowchart (purple box represents the step of identifying the initial MEQ and and the green box represents the step of relocation of the MEQ hypocenter)

Geiger Method

Geiger Method is one of the inversion methods for determining the initial MEQ hypocenter (Madrinovella, 2012; Midzi et al., 2020; Nishi, 2005). The Geiger method is applied in determining a single hypocenter by using residual time data, where this data is the difference between the observation time and the calculation time (Chen & Huang, 2018; Riziq Maulana et al., 2019; Utama et al., 2021). The GAD method control is the seismotectonic state suitability of the research as indicated by the RMSE value. Smaller RMSE value (close to zero), generally indicates the better or more accurate MEQ hypocenter determination. The RMSE value in this

method is obtained from the residual observation time and the earthquake occurrence time.

Coupled Velocity-Hypocenter Method

Coupled Velocity Hypocenter (CVH) Inversion Method is an earthquake hypocenter testing method with simultaneous station correction using the Geiger method (Kianimehr et al., 2018; Kissling, 1995, 1998; Sevilla et al., 2020; Utama, Garini, & Indriani, 2022). The CVH method is a travel time inversion method, where the intended travel time is the difference between the seismic wave arrival time and the earthquake event time. The CVH inversion results are the earthquake hypocenter locations in the form of coordinates, depth, RMSE value, and azimuth gap.

Result and Discussion

The data has been processed with variations in the different station distributions using the CVH and GAD inversion methods, showing the MEQ hypocenters distribution results from 8 seismic stations. It is known that many MEQ hypocenter locations distribution is identified in zones that are relatively close to the presence of seismic stations. This is shown in Figure 2, where the MEQ hypocenters determined by the 8 seismic stations are clustered in the zone around the 8 stations, and the distance between hypocenters are increasingly stretched as the hypocenter are further away from the stations. This is supported by the hypocenter distribution pattern, when viewed from the top as shown in Figure 3. The hypocenter cluster has a high concentration in the zone relatively close to the seismic station and is increasingly tenuous when the hypocenter position is relatively far from the station.

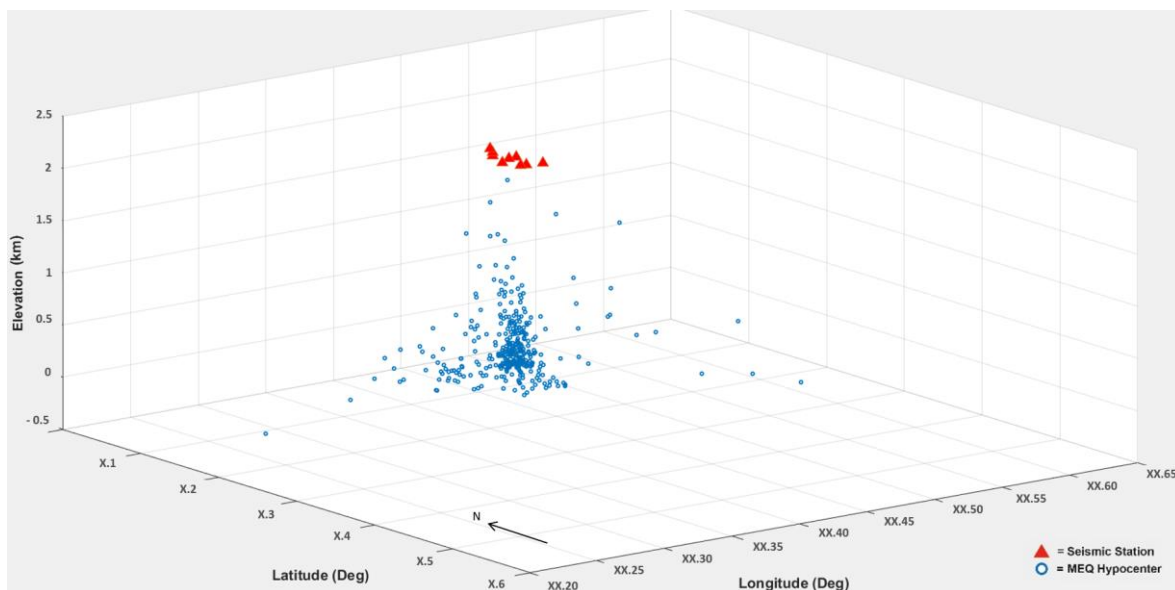


Figure 2. CVH inversion data plot results for 8 stations in 3D model

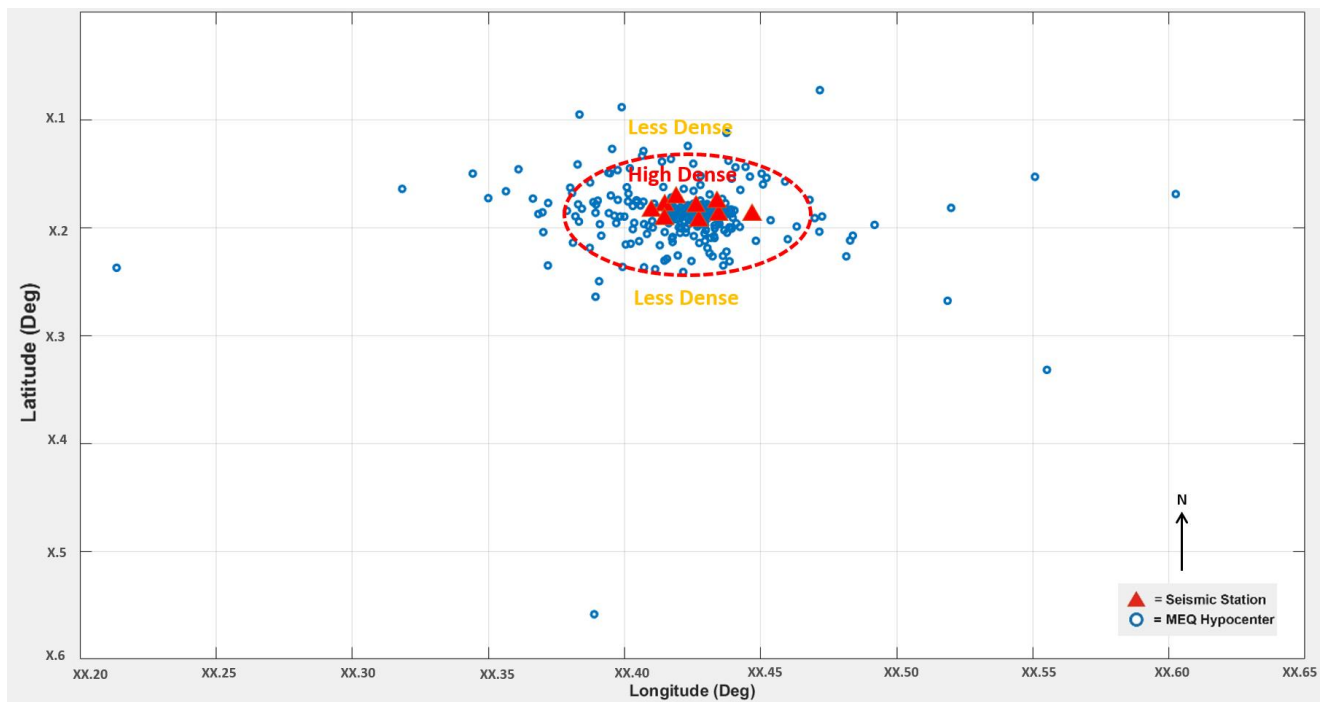


Figure 3. CVH inversion data plot results for 8 stations in bird view

Hypocenter distribution pattern from the cross-section is also interesting to discuss. In the latitude coordinate section as shown in Figure 4, it is known that the hypocenter point density has a high concentration just below the station located at a depth of 1.5 to 2 kilometers below the seismic station. The relatively dense hypocenter point has no effect spatially vertically, but is affected spatially horizontally. This is also shown in Figure 5, which displays latitude coordinate incisions, where the MEQ hypocenters distribution pattern forms a "triangle-shape", showing the depth effect does not significant on the hypocenter location identification, but the hypocenter distribution in horizontal axis gave more significant effect on the hypocenter point identification. In other words, if the MEQ hypocenter is closer to the seismic station, there are more hypocenters will be detected by the seismic station.

After knowing the relationship between MEQ hypocenters distribution and the seismic station location, to determine the seismic stations number effect on the MEQ hypocenters distribution, seismic stations will be divided, which originally consisted of 8 stations into 2 groups of stations, each group consist of 4 stations. The data below station position, with horizontally spaced intervals, that are relatively close to the station zone, tends to lower RMSE value. This support by analysis results in the previous section, where the hypocenter point will be "more identified" in the relatively close to the station zone horizontally. It is known that the depth factor does not have a significant influence on the MEQ hypocenters distribution identified by seismic stations.

Average RMSE values obtained from stations 4A and 4B were 0.159 and 0.105, respectively. This result is lower than the result from 8 stations simultaneously, which has an average RMSE value of 0.26. It is known that the less number of seismic stations in MEQ hypocenter measurement impacts the lower RMSE value. This is inseparable from the station's location influence and the amount of hypocenter data used, where is farther hypocenter from the seismic station and the more hypocenter data calculated will potentially produce a larger RMSE value.

The MEQ hypocenter results plot based on the grouping of 4 stations produce a relatively wide distribution of hypocenter points compared to the inversion results at 8 stations. Figure 6 shows the data distribution from the CVH inversion results from station groups 4A and 4B. The hypocenter points relatively high concentration emergence in the zone that relatively close to the station location. This is supported by the top view plot results shown in Figure 7, where the fewer number of stations (4 seismic stations) results in a relatively more spread-out hypocenter points distribution, although the denser zones of hypocenter points are still in a zone that is relatively close to the station locations.

The MEQ hypocenter points distribution in the 4 station grouping is more spread-out as a small number effect of seismic stations, it turns out that it does not only occur horizontally, but also vertically. This is shown in Figure 8 and Figure 9, where the latitude and longitude plots results will identify the MEQ hypocenter points distribution which are increasingly spread out as there are fewer number of seismic stations. Even though there

is also a distribution of hypocenter points that spread more vertically, it is still possible to identify dense zones from the MEQ hypocenter point, where the location of the dense zone is relatively right below the seismic station. This is related to the azimuth gap where the average value of stations 4A and 4B is 283 and 267, whereas when using 8 stations the azimuth gap average value is 222. High azimuth gap value in the data analysis results is caused by the stations distribution which are only on one side, not scattered around the earthquake epicenter. The MEQ hypocenter plot results from the

CVH inversion results with the grouping of 4 stations show a tendency towards a certain direction. The direction orientation formed from the MEQ hypocenter plot is influenced by the location of the seismic station. Station 4A which is relatively in the western part, records more hypocenters in the western cross-sectional area, while station 4B which is relatively in the eastern part records more hypocenters in the eastern cross-sectional area. This shows MEQ hypocenter distribution pattern indication from data which tends to follow the location of the seismic measurement station.

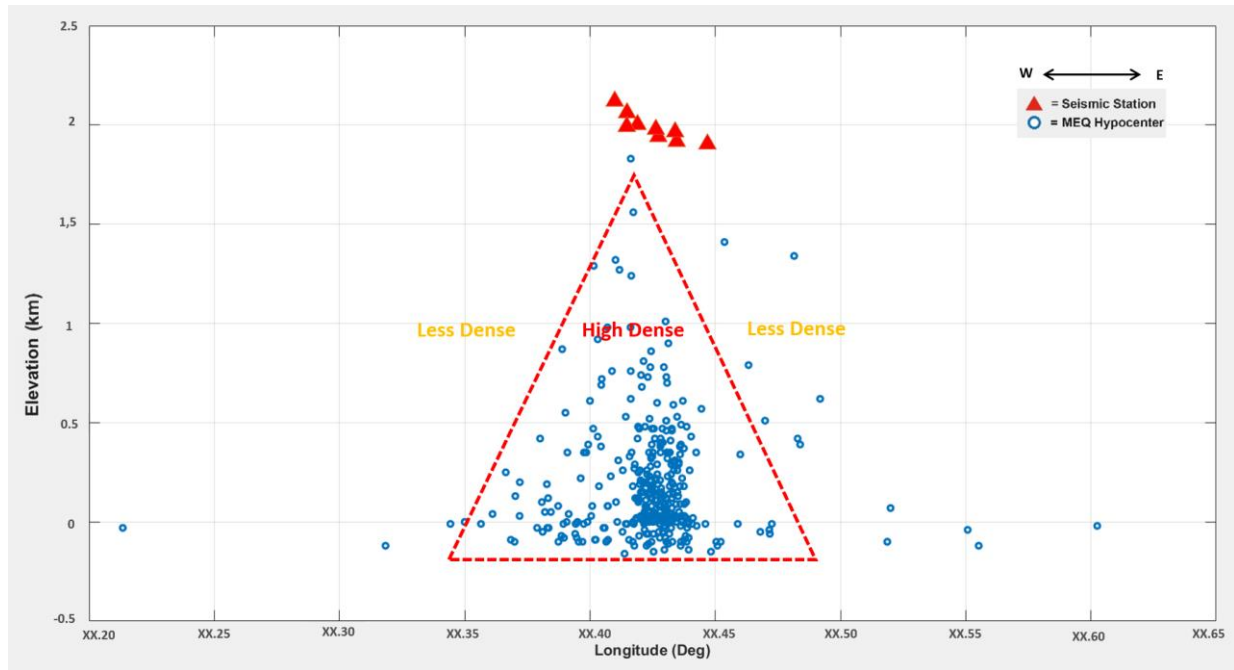


Figure 4. CVH inversion data plot results for 8 stations in longitude coordinate section

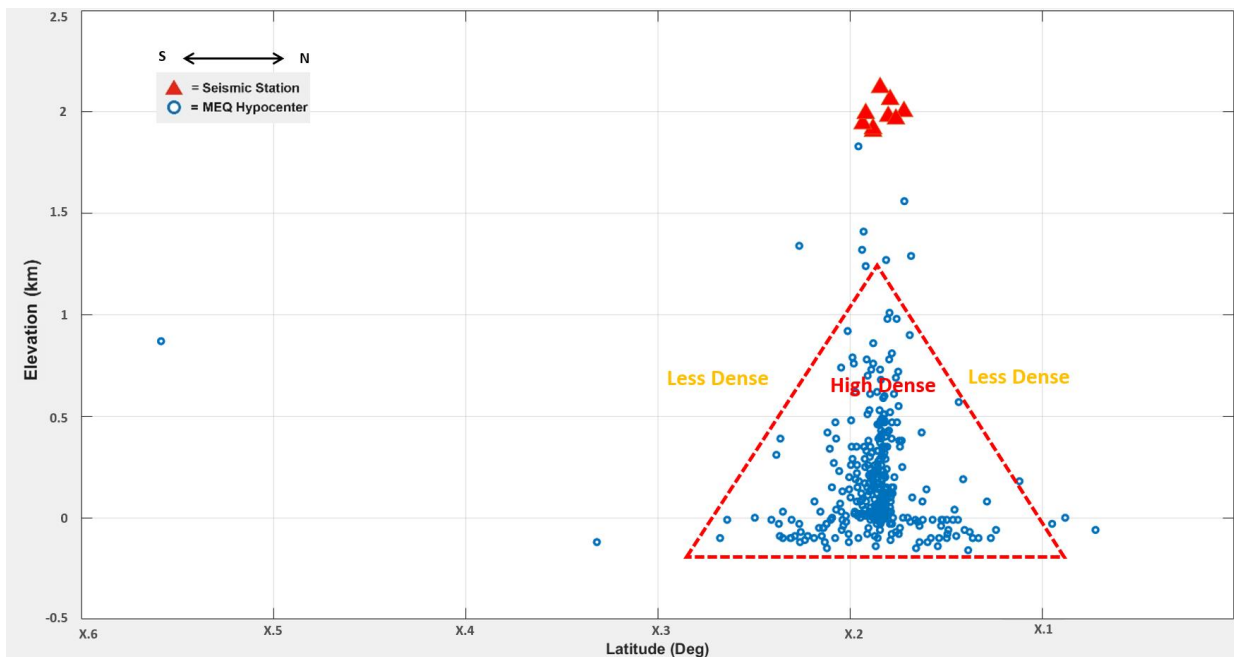


Figure 5. CVH inversion data plot results for 8 stations in latitude coordinate section

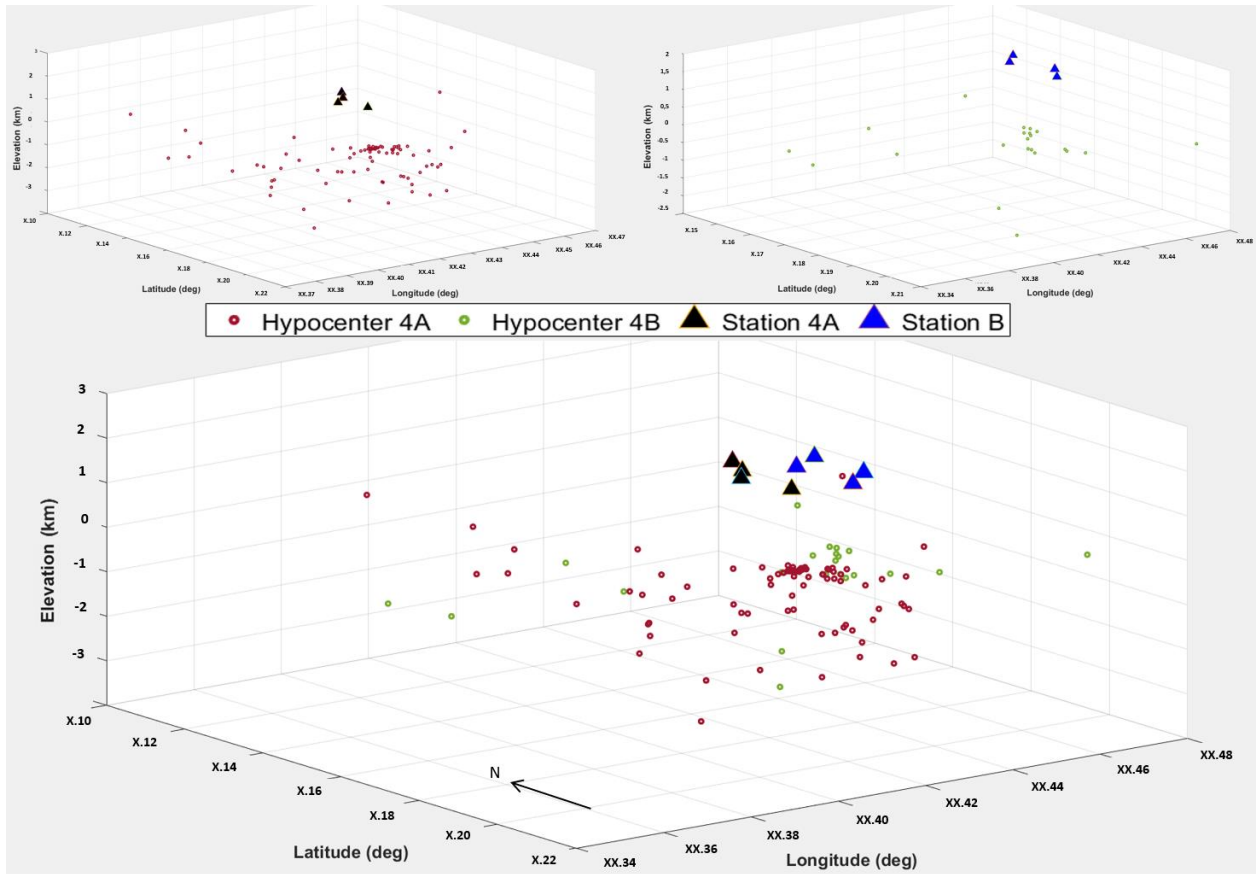


Figure 6. CVH inversion data plot results for 4 stations grouping in 3D model

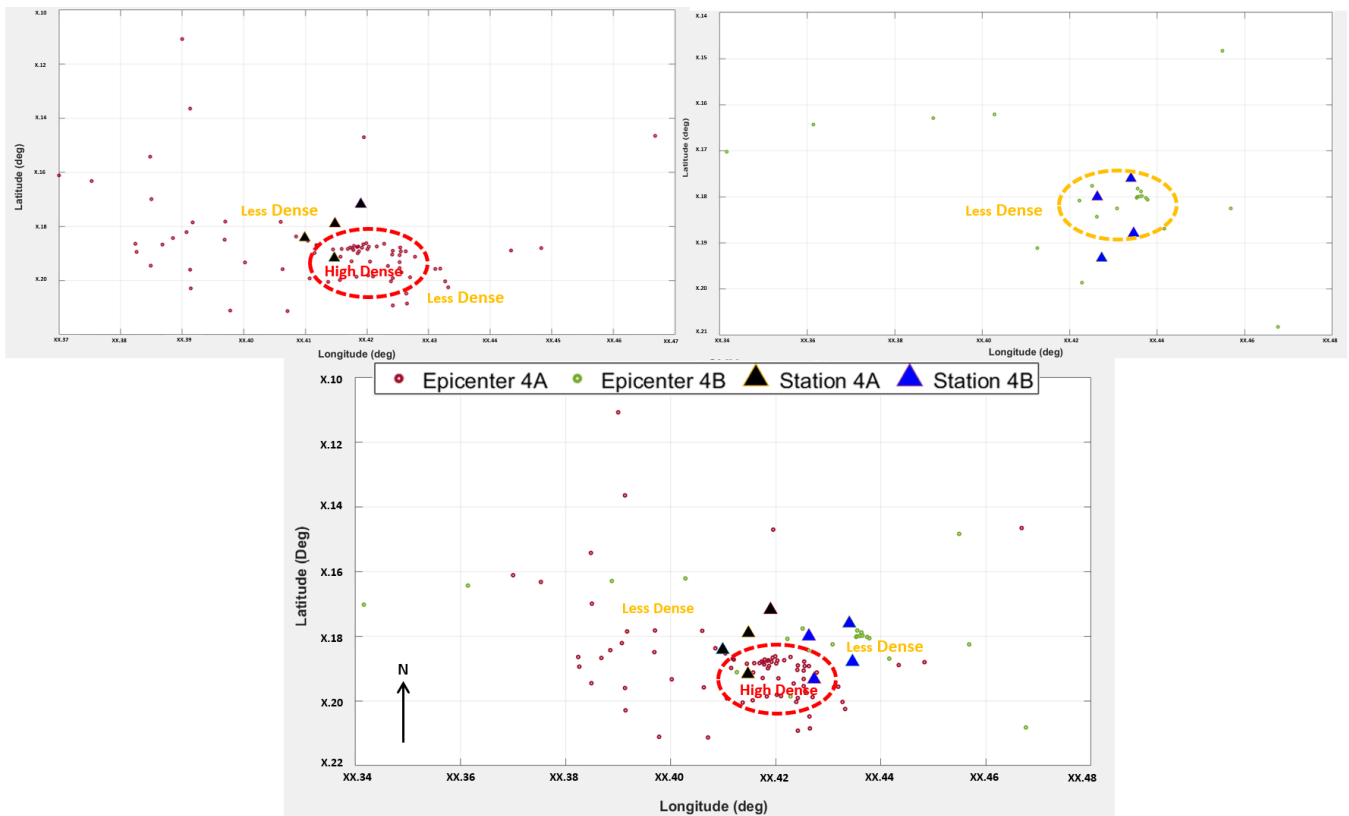


Figure 7. CVH inversion data plot results for 4 stations grouping in bird view

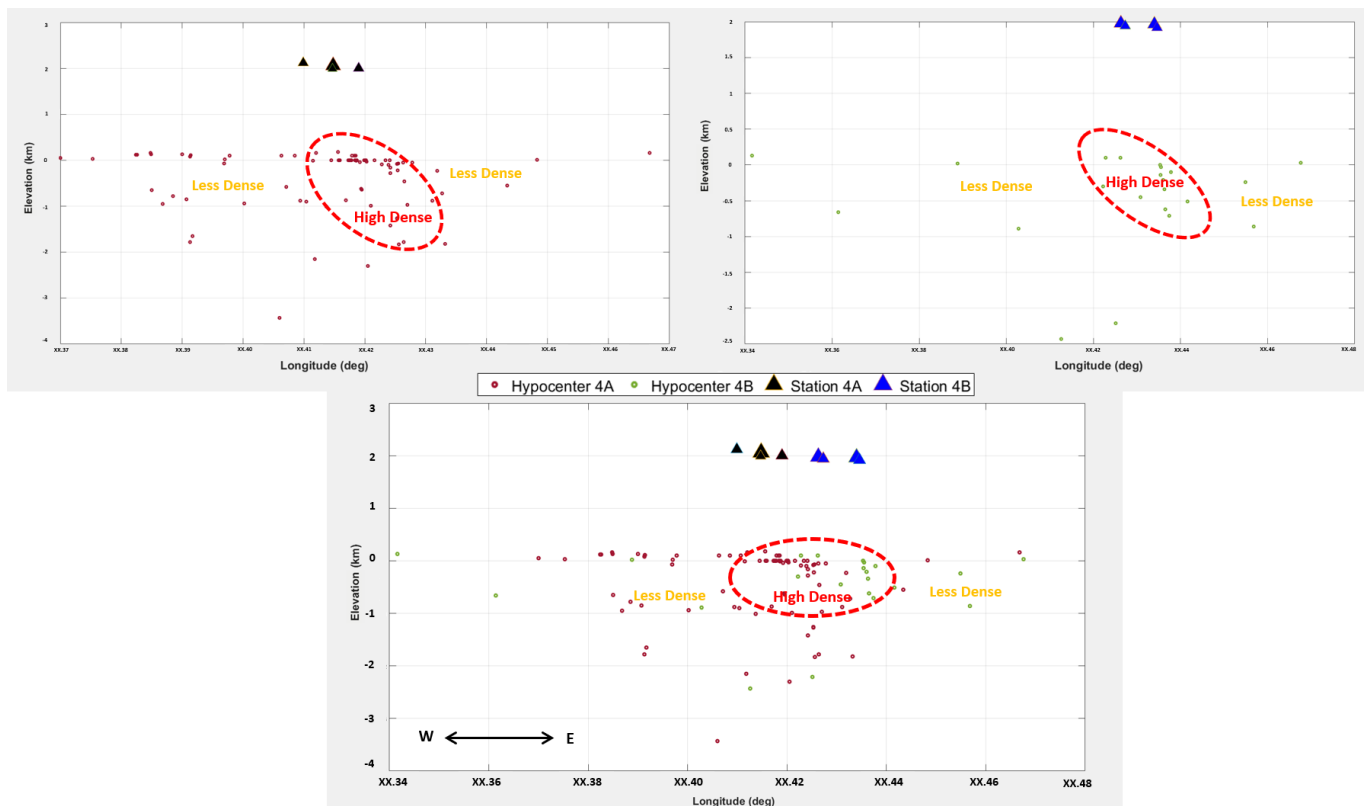


Figure 8. CVH inversion data plot results for 4 stations grouping in longitude coordinate system

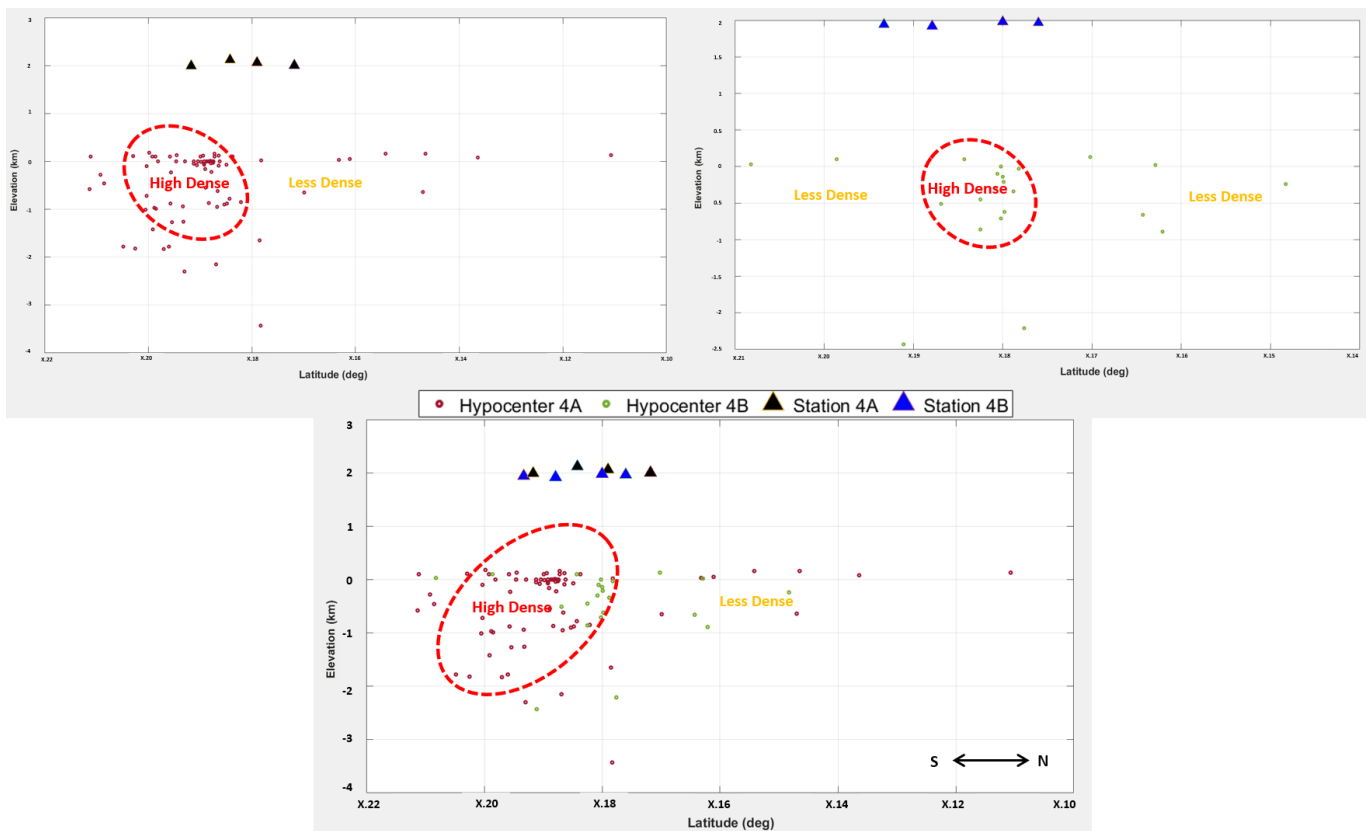


Figure 9. CVH inversion data plot results for 4 stations grouping in latitude coordinate system

The approach that can be taken as an effort to minimize the MEQ location bias is to increase the number of seismic stations 360° around the earthquake event in the study area, so the azimuth gap can be below 120° (Bidang, 2020; Bondár et al., 2004; Fahntalia & Madlazim, 2017). Even though each method is given a minimum limit for the use of seismic stations, it would be nice to increase the number of seismic stations beyond the minimum standard. If this is not possible, then with the number of existing stations, an initial study is carried out regarding the history of the existing reservoirs so that the placement of seismic stations can be efficient and can produce a good picture of the subsurface of the reservoir.

Conclusion

MEQ hypocenter data processing using the CVH and GAD methods results show a different micro-earthquake hypocenter locations distribution for each type of station distribution. MEQ hypocenter distribution was mostly identified in zones relatively close to seismic stations, with the intervals between hypocenters stretching further as the hypocenters move away from the station. Experimental results show that in determining the location of the MEQ, the geometry and number of seismic stations played an important role. Increasing the number of stations with relatively long distances can result in less accurate locations of MEQ, error and bias in determining the location of MEQ will be greater when the azimuth gap value is greater. This is shown by the distribution of MEQ that are more spread out in variations 4A and 4B (4 seismic stations) compared to the distribution of MEQ hypocenters using data from 8 seismic stations. The azimuth gap variations of stations 4A and 4B are 283° and 267°, and 8 stations have a value of 222°. The large value of the azimuth gap is due to the distribution of stations only on one side so that there are horizontal angles that are not covered by seismic stations. An approach that can be taken to minimize MEQ location bias is to increase the number of 360° seismic stations around earthquake events in the study area, so that the azimuth gap can be below 120°.

Acknowledgments

The authors would like to thank Institut Teknologi Sepuluh Nopember for the support provided for this research under "Penelitian Dasar Lanjutan" scheme with contract No. 008/E5/PG.02.00.PT/2022 dated March 16, 2022.

Author Contributions

The author's contributions include Widya Utama, Sherly A. Garini, Dwa D. Warnana, and Wien Lestari: focus on methodology and review on writing, and so on. Sherly A. Garini, Dhea P. N. Putra, and Merry C. Hutapea: data

collection, data processing, data analysis, and writing the original manuscript.

Funding

This research was funded by Institut Teknologi Sepuluh Nopember under "Penelitian Dasar Lanjutan" scheme with contract No. 008/E5/PG.02.00.PT/2022 dated March 16, 2022.

Conflicts of Interest

The authors declare no conflict of interest regarding the publication of this paper.

References

- Anderson, A., & Rezaie, B. (2019). Geothermal technology: Trends and potential role in a sustainable future. *Applied Energy*, 248(1), 18–34. <https://doi.org/10.1016/j.apenergy.2019.04.102>.
- Aneke, M. C., & Menkiti, M. C. (2016). Geothermal: History, Classification, and Utilization for Power Generation. In *Alternative Energy and Shale Gas Encyclopedia*, 251–264. <https://doi.org/10.1002/9781119066354.ch26>
- Barbosa, N. D., Solazzi, S. G., & Lupi, M. (2020). Seismically Induced Unclogging in Fluid-Saturated Faults. *Journal of Geophysical Research: Solid Earth*, 125(8), 1–20. <https://doi.org/10.1029/2020JB020152>.
- Bertani, R. (2016). Geothermal power generation in the world 2010-2014 update report. *Geothermics*, 60, 31–43. <https://doi.org/10.1016/j.geothermics.2015.11.003>.
- Bidang, A. W. (2020). Pengaruh Penambahan Stasiun-Stasiun Seismik (STPI, TSPI, dan IWPI) Terhadap Analisa Penentuan Parameter Gempa Bumi Studi Kasus Gempa Bumi Di Wilayah Papua Barat Tahun 2019 – 2020. *Biolearning Journal*, 34(1), 1–17. <https://doi.org/10.15233/gfz.2017.34.5>.
- Bondár, I., Myers, S. C., Engdahl, E. R., & Bergman, E. A. (2004). Epicentre accuracy based on seismic network criteria. *Geophysical Journal International*, 156(3), 483–496. <https://doi.org/10.1111/j.1365-246X.2004.02070.x>
- Chen, Y., & Huang, L. (2018). Microearthquake Hypocenter-Location and Focal-Mechanism Inversions for the EGS Collab Project: A Synthetic Study. *Pangea.Stanford.Edu*, 1–6. Retrieved from <https://pangea.stanford.edu/ERE/pdf/IGastandard/SGW/2018/Chen.pdf>
- Cheng, A. (2022). President's Page: Geothermal energy: Current and future. *Leading Edge*, 41(9), 588–589. <https://doi.org/10.1190/tle41090588.1>.
- Fahntalia, C. P., & Madlazim. (2017). Pengaruh Jumlah Stasiun Seismik Terhadap Hasil Estimasi Centroid Moment Tensor Gempa Bumi. *Jurnal Inovasi Fisika*

- Indonesia (IFI)*, 6(3), 1–5. Retrieved from <https://ejournal.unesa.ac.id/index.php/inovasi-fisika-indonesia/article/view/19899>
- Farhan, F., Syamsu Rosid, M., Riziq Maulana, M., & Iskandar, C. (2019). The First Study of 3-D Seismic Velocities Tomography Inversion to Delineate Reservoir Steam Zone at “fR” Geothermal Field West Java. *Journal of Physics: Conference Series*, 1351(1). <https://doi.org/10.1088/17426596/1351/1/012049>.
- Firdaus Al Hakim, M., Sule, R., & Hendriyana, A. (2019). Seismicity Analysis and Velocity Structure of the Two-Phase Geothermal Field in West Java, Indonesia: Preliminary Result. *IOP Conference Series: Earth and Environmental Science*, 318(1). <https://doi.org/10.1088/17551315/318/1/012039>
- Geoffroy, L., Dorbath, C., Ágústsson, K., Kristjánssdóttir, S., Flóvenz, Ó. G., Doubre, C., Gudmundsson, Ó., Barreyre, T., Bazin, S., & Franco, A. (2022). Hydrothermal fluid flow triggered by an earthquake in Iceland. *Communications Earth and Environment*, 3(1). <https://doi.org/10.1038/s43247-022-00382-0>.
- Halim, G. R., Utama, W., & Mariyanto, M. (2020). Uji Lokasi Hiposenter Mikro-Earthquake (Meq) dengan Metode Inversi Simulated Annealing pada Lapangan Panas Bumi “XX.” *Jurnal Geosaintek*, 6(2), 71. <https://doi.org/10.12962/j25023659.v6i2.6548>.
- Huang, W., Wang, R., & Chen, Y. (2018). Regularized non-stationary morphological reconstruction algorithm for weak signal detection in microseismic monitoring: Methodology. *Geophysical Journal International*, 213(2), 1189–1211. <https://doi.org/10.1093/gji/ggy054>
- Intani, R. G., Golla, G. U., Syaffitri, Y., Paramitasari, H. M., Nordquist, G. A., Nelson, C., Ginanjar, Giri, G. K. D. S., & Sugandhi, A. (2020). Improving the conceptual understanding of the Darajat Geothermal Field. *Geothermics*, 83(September). <https://doi.org/10.1016/j.geothermics.2019.101716>
- Jouhara, H., Żabnieńska-Góra, A., Khordehgah, N., Ahmad, D., & Lipinski, T. (2020). Latent thermal energy storage technologies and applications: A review. *International Journal of Thermofluids*, 5–6, 100039. <https://doi.org/10.1016/j.ijft.2020.100039>.
- Karasözen, E., & Karasözen, B. (2020). Earthquake location methods. *GEM - International Journal on Geomathematics*, 11(1). <https://doi.org/10.1007/s13137-020-00149-9>.
- Kato, A., Sakai, S., Matsumoto, S., & Iio, Y. (2021). Conjugate faulting and structural complexity on the young fault system associated with the 2000 Tottori earthquake. *Communications Earth and Environment*, 2(1), 1–10. <https://doi.org/10.1038/s43247-020-00086-3>.
- Kianimehr, H., Kissling, E., Yaminifard, F., & Tatar, M. (2018). Regional minimum 1-D P-wave velocity model for a new seismicity catalogue with precise and consistent earthquake locations in southern Iran. *Journal of Seismology*, 22(6), 1529–1547. <https://doi.org/10.1007/s10950-018-9783-4>.
- Kissling, E. (1998). Geotomography with Local Earthquake Data. *Journal of Chemical Information and Modeling*, 26(4), 6. <https://doi.org/10.1017/CBO9781107415324.004>
- Kissling, E., Kradolfer, U., & Murer, H. (1995). *Program VELEST USER'S GUIDE - Short Introduction*. In Institute of Geophysics, ETH Zurich
- Küperkoch, L., Olbert, K., & Meier, T. (2018). Long-term monitoring of induced seismicity at the Insheim geothermal site, Germany. *Bulletin of the Seismological Society of America*, 108(6), 3668–3683. <https://doi.org/10.1785/0120170365>.
- Lund, J. W., Freeston, D. H., & Boyd, T. L. (2011). Direct utilization of geothermal energy 2010 worldwide review. *Geothermics*, 40(3), 159–180. <https://doi.org/10.1016/j.geothermics.2011.07.04>.
- Madrinovella, I. (2012). Studi Penentuan dan relokasi Hiposenter Serta Mekanisme Fokus Gempa Mikrodi Sekitar Cekungan Bandung. *Jurnal Geofisika*, 13(2), 80–88.
- Mahwa, J., Li, D. jiang, Ping, J. hua, Leng, W., Tang, J. bo, & Shao, D. yun. (2022). Mapping the spatial distribution of fossil geothermal manifestations and assessment of geothermal potential of the Tangyin rift, Southeast of Taihang Mountain in China. *Journal of Mountain Science*, 19(8), 2241–2259. <https://doi.org/10.1007/s11629-022-7329-2>.
- Midzi, V., Pule, T., Manzunzu, B., Mulabisana, T., Zulu, B. S., & Myendeki, S. (2020). Improved earthquake location in the gold mining regions of south africa using new velocity models. *South African Journal of Geology*, 123(1), 35–58. <https://doi.org/10.25131/sajg.123.0008>.
- Moya, D., Aldás, C., & Kaparaju, P. (2018). Geothermal energy: Power plant technology and direct heat applications. In *Renewable and Sustainable Energy Reviews*, 94, 889–901. <https://doi.org/10.1016/j.rser.2018.06.047>.
- Nishi, K. (2005). *Hypocenter Calculation Software GAD (Geiger's method with Adaptive Damping) GAD Manual Guide*.
- Østergaard, P. A., Duic, N., Noorollahi, Y., Mikulcic, H., & Kalogirou, S. (2020). Sustainable development using renewable energy technology. *Renewable Energy*, 146, 2430–2437.

- <https://doi.org/10.1016/j.renene.2019.08.094>.
- Pennington, C. N., Chang, H., Rubinstein, J. L., Abercrombie, R. E., Nakata, N., Uchide, T., & Cochran, E. S. (2022). Quantifying the Sensitivity of Microearthquake Slip Inversions to Station Distribution Using a Dense Nodal Array. *Bulletin of the Seismological Society of America*, 112(3), 1252–1270. <https://doi.org/10.1785/0120210279>.
- Rahmaningtyas, N. I., Utama, W., & Lestari, W. (2020). Analisis Sumber Gempa Mikro Melalui Distribusi Lokasi Hiposenter Menggunakan Metode Double Difference pada Lapangan Panas Bumi "X." *Jurnal Geosaintek*, 6(1), 33. <https://doi.org/10.12962/j25023659.v6i1.6549>
- Riziq Maulana, M., Syamsu Rosid, M., Farhan, F., & Iskandar, C. (2019). Identification of Fracture Density and Orientation at "r" Geothermal Field Using Shear Wave Splitting Microearthquake Method. *Journal of Physics: Conference Series*, 1351(1). <https://doi.org/10.1088/17426596/1351/1/012050>.
- Sevilla, W. I., Jumawan, L. A., Clarito, C. J., Quintia, M. A., Dominguiano, A. A., & Solidum, R. U. (2020). Improved 1D velocity model and deep long-period earthquakes in Kanlaon Volcano, Philippines: Implications for its magmatic system. *Journal of Volcanology and Geothermal Research*, 393, 106793. <https://doi.org/10.1016/j.jvolgeores.2020.106793>.
- Sicking, C., & Malin, P. (2019). Fracture seismic: Mapping subsurface connectivity. *Geosciences (Switzerland)*, 9(12). <https://doi.org/10.3390/geosciences9120508>.
- Siddiq, N. A., Chong, W. Y., Pramono, Y. H., Muntini, M. S., Asnawi, A., & Ahmad, H. (2020). All-Optical Humidity Sensor Using SnO₂ Nanoparticle Drop Coated on Straight Channel Optical Waveguide. *Photonic Sensors*, 10(2), 123–133. <https://doi.org/10.1007/s13320-019-0563-8>.
- Toledo, T., Jousset, P., Maurer, H., & Krawczyk, C. (2020). Optimized experimental network design for earthquake location problems: Applications to geothermal and volcanic field seismic networks. *Journal of Volcanology and Geothermal Research*, 391(3). <https://doi.org/10.1016/j.jvolgeores.2018.08.011>.
- Utama, W., Ardhya, S., Anggita, V., Lestari, W., & Desa, D. (2022). Delay Time (δt) and Polarization Direction (φ) Analysis Based on Shear Wave Splitting (SWS) Method. *International Journal on Advanced Science, Engineering and Information Technology*, 12(5), 2075–2082. <https://doi.org/10.18517/ijaseit.12.5.13100>
- Utama, W., & Garini, S. A. (2022). Hypocenter Determination and Estimation 1-D Velocity Models Using Coupled Velocity-Hypocenter Method. *International Journal on Advanced Science, Engineering and Information Technology*, 12(3), 892–898. <https://doi.org/10.18517/ijaseit.12.3.12488>.
- Utama, W., Garini, S. A., & Indriani, R. F. (2022). Distribution Analysis of Micro-earthquakes in Geothermal Areas by using Coupled Velocity-Hypocenter and Double Difference Methods. *In Review : Jurnal Teknologi*, 1(2016), 1–5.
- Utama, W., Garini, S. A., & Lansia, F. S. (2022). The Effect of Picking Uncertainty Window Interval (ΔT_p) on Hypocenter Micro-Earthquake (MEQ) Location using Geiger Method. *Journal of Marine-Earth Science Technology*, 3(1), 5–10. <https://doi.org/10.12962/j27745449.v3i1.439>
- Utama, W., Warnana, D. D., & Garini, S. A. (2021). Identification of Micro-Earthquake Hypocentre using Geiger and Coupled Velocity-Hypocentre Methods. *International Journal on Advanced Science, Engineering and Information Technology*, 11(1), 350–355. <https://doi.org/10.18517/ijaseit.11.1.10589>
- Wildan Perdana, M., Mendrofa, D., Lubis, T., Nordquist, G., Energy, S., & Salak, G. (2020). Microearthquake Monitoring at Salak Geothermal Field, Indonesia: Application of Tomography Inversion and Its Impact on Reservoir Characterization Precision Gravity and Leveling for Salak Geothermal Field Monitoring View project Muhamad Wildan Perdana Star. *World Geothermal Congress, May*. Retrieved from <https://www.researchgate.net/publication/356283797>.
- Yang, Z., Yehya, A., Iwalewa, T. M., & Rice, J. R. (2021). Effect of Permeability Evolution in Fault Damage Zones on Earthquake Recurrence. *Journal of Geophysical Research: Solid Earth*, 126(9), 1–17. <https://doi.org/10.1029/2021JB021787>.
- Zaini, N., Yanis, M., Abdullah, F., Van Der Meer, F., & Aufaristama, M. (2022). Exploring the geothermal potential of Peut Sageo volcano using Landsat 8 OLI/TIRS images. *Geothermics*, 105(June), 102499. <https://doi.org/10.1016/j.geothermics.2022.102499>.
- Zhang, M., Ellsworth, W. L., & Beroza, G. C. (2019). Rapid Earthquake Association and Location. *Seismological Research Letters*, 90(6), 2276–2284. <https://doi.org/10.1785/0220190052>.