

# Detection of hydrological mass variations by means of an inverse tesseroid approach

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Numerical results of the closed-loop simulation

## Introduction

storage values (TWS), e.g. Awange et al. (2009). values (Heck and Seitz, 2008).

Modern satellite missions like GRACE (Gravity contribution, an approach based on tesseroidal Recovery and Climate Experiment) provide mass bodies is proposed (Heck and Seitz, monthly global geopotential models allowing the 2007). Tesseroids are bounded by geographical analysis of the time-variable part of the Earth's grid lines and surfaces of constant heights and gravity field. From a hydrological point of view, thus directly linked to the curvature of the Earth. the monthly changes are mainly caused by Therefore, tesseroid formulas are introduced as variations of the Earth's water storage. The observation equations in a least-squares usual procedure for the detection of such adjustment, in which gravity data, e.g. derived hydrological mass variations is based on the from GRACE, are utilized as observations and elastic Love number theory, in which global the heights of particular tesseroids are the spherical harmonics are used to derive geoid unknowns. Using a constant density value of anomalies that can be converted into water 1000 kg/m3, the solved unknown tesseroid column heights respectively terrestrial water heights can be associated with the desired TWS

Alternative procedures in the space domain Taking the Australian continent as an result in an inverse problem of forward example, results from a realistic closed-loop modelling. This procedure is particularly suitable simulation are presented, showing the for regional applications. In contrast to the often possibilities of the suggested tesseroid method used point-mass approximation, in this compared to a point mass approximation.

## **Closed-loop simulation**

Starting with 1°x1° gridded TWS values of the In the second step, the generated potential year 2010 over Australia (Dijk et al., 2011), monthly TWS anomalies are constituted by subtracting the annual average.

In the first step, these TWS anomalies are associated with heights of tesseroids that are used in forward modelling to determine a 1°x1° grid of consistent potential values at satellite altitude. As the volume integral occurring in the tesseroid formula cannot be solved analytically. a Taylor series expansion of the integral kernel is carried out (Heck and Seitz, 2007).

values are introduced in an iterative leastsquares adjustment in order to reconstruct the TWS information by inverse modelling. In the first iteration, the linear part of the observation equation is used to get approximate parameters which can be associated with a point mass. These parameters then serve as approximate values for the linearized observation equation. After each iteration step, the estimated parameters are compared to the original TWS anomalies producing residual values.





# **Conclusions and Outlook**

- Closed-loop simulation shows a good behaviour for the developed inverse tesseroid approach.
- Compared to point mass approximation (1. iteration), the tesseroid method shows a significant reduction of the residual TWS values by three orders of magnitude.
- The magnitude of the TWS residuals highly depend on the reference height h of the potential grid.
- For the desired application of GRACE (h = 450 km) two iteration steps already show a sufficient accuracy at sub-millimeter level.
- TWS residuals for different seasons show similar grid effects which will be investigated further.
- The stability of the inverse tesseroid method in the case of noisy observations will be analysed
- Tesseroid method will be applied to real measured GRACE data and compared to the results of alternative approaches in the frequency domain.

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### Acknowledgements:

The authors would like to thank A. van Dijk, CSIRO Land and Water, for kindly providing TWS values modeled by the Australian Water Resources Assessment (AWRA) system and E. Forootan, University of Bonn, for making contact to CSIRO

