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Changes in the relative importance of biogenic isoprene and Soil NO_x emissions on ozone concentrati...

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Key Points:

- Following anthropogenic nitrogen oxides (NO_x) reductions, O₃ sensitivity to biogenic isoprene is decreasing in many US O₃-nonattainment areas
- Soil NO_x dominates isoprene in interannual tendencies of O₃ in nonattainment areas of the Great Lakes, Ohio River Valley, and Northeast (NEA)
- Emerging importance of soil NO_x emissions implies potentially unexpected variability in high O₃ where volatile organic compound-limitation previously dominated

Supporting Information:

Supporting Information may be found in the online version of this article.

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Changes in the Relative Importance of Biogenic Isoprene and Soil NO_x Emissions on Ozone Concentrations in Nonattainment Areas of the United States

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Abstract Reductions in anthropogenic emissions have drawn increasing attention to the role of the biosphere in O₃ production chemistry in U.S. cities. We report the results of chemical transport model sensitivity simulations exploring the relative impacts of biogenic isoprene and soil nitrogen oxides (NO_x) emissions on O₃ and its temporal variability. We compare scenarios with high and low anthropogenic NO_x emissions representing the reductions that have occurred in recent decades. As expected, summertime O₃ concentrations become less sensitive to perturbations in biogenic isoprene emissions as anthropogenic NO_x emissions decline. However, we demonstrate for the first time that across policy relevant O₃ nonattainment areas of the United States, O₃ becomes more sensitive to perturbations in soil NO_x emissions than to identical perturbations in isoprene emissions. We show that interannual variability in soil NO_x emissions may now have larger impacts on interannual O₃ variability than isoprene emissions in many areas where the latter would have dominated in the recent past.

Plain Language Summary Ozone (O₃) is a criteria air pollutant that continues to pose a threat to more than one hundred million Americans each year, despite progress in regulating precursor emissions. In many O₃-polluted areas, the role of natural emissions of isoprene in the production of ground-level O₃ has been well recognized, but this chemistry depends strongly on local anthropogenic emissions which have been changing rapidly. We use an updated estimate of anthropogenic emissions to demonstrate that many areas that remain in nonattainment of the federally mandated O₃ standard are now much less sensitive to natural isoprene emissions, with biogenic nitrogen oxide emissions from soils becoming more and more important. The role of these soil emissions on O₃ in nonattainment areas has not been well characterized, but, as we show here, this will become increasingly necessary for good O₃ policy in nonattainment areas.

1. Introduction

Despite widespread improvements in air quality across the United States, more than 120 million Americans continue to live in regions that do not meet the current health-based ground-level ozone (O₃) standards (U.S. Environmental Protection Agency, 2019b). Known as “nonattainment areas,” these areas have O₃ design values (defined as a 3-year average of the annual fourth-highest maximum 8 hr O₃ mixing ratio measured at an air quality monitoring site) that exceed 70 ppb (U.S. Environmental Protection Agency, 2019a). As anthropogenic precursor emissions have declined, we require increasing detail regarding the role of natural variability and background O₃ in local air quality management (Fiore et al., 2014; Guo et al., 2018; Jaffe et al., 2018; Lin et al., 2017; Parrish et al., 2017). This includes contributions from biogenic precursor emissions, such as volatile organic compounds (VOCs) from vegetation and nitrogen oxides (NO_x) from soils. While some portion of these emissions may not be considered “natural” (e.g., NO_x emissions from managed agricultural soils), their emission rates depend strongly on natural environmental variability.

Because O₃ chemistry in urban core areas has historically been largely VOC-limited (i.e., O₃ production is positively proportional to changes in organic compound emissions), past focus has been on quantifying the influence of biogenic isoprene emissions in particular (Bell & Ellis, 2004; Chameides et al., 1988; Hakami et al., 2006; Kleinman et al., 2005; Li et al., 2007; Tao et al., 2003; Zare et al., 2014). With wide reductions in NO_x emissions across the United States (Duncan et al., 2016; Hidy & Blanchard, 2015; Parrish et al., 2011), satellite- and ground-based monitoring analyses demonstrate that O₃ production in urban areas of the country is transitioning from mostly VOC-limited to increasingly NO_x-limited (Baidar et al., 2015; Blanchard & Hidy, 2018; He

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et al., 2019; Jin et al., 2017; Pusede & Cohen, 2012; Pusede et al., 2015). Recent observational evidence has subsequently supported the growing importance of soil NO_x emissions on O₃ levels in some areas (Almaraz et al., 2018; Oikawa et al., 2015; Romer et al., 2018; Trousdell et al., 2019), but this work has been limited to short-term field campaigns in specific regions (e.g., California and the Southeast United States), without attention on the potential impacts across all U.S. nonattainment areas.

Biogenic emissions exhibit large temporal variability, so that their importance and relative contributions vary day-to-day and year-to-year. Both have underlying sensitivity to some overlapping environmental variables but could potentially drive changes in O₃ concentrations in different ways that would benefit from being understood fully under changing anthropogenic emissions. The influence of the biosphere on O₃ air quality during extreme climatological events or during the worst air quality days has been confirmed (Demetillo et al., 2019; Guo et al., 2018; Wang et al., 2017; Zhang & Wang, 2016). For example, the enhancement in O₃ from biogenic VOC emissions on the 10 highest O₃ days in the United States over 2004–2012 was estimated to be 1–7 ppb higher than their contribution to O₃ under more average conditions (Guo et al., 2018). Romer et al. (2018) estimate that soil NO_x emissions are responsible for almost half of the 2.3 ppb °C⁻¹ increase in O₃ production rates with ambient temperatures in the southeast United States. This variability in biogenic emissions is often invoked as a mechanism behind well-documented correlations between O₃ concentrations and ambient temperature (Lu et al., 2019; Pusede et al., 2014, 2015; Rasmussen et al., 2012; Romer et al., 2018; Steiner et al., 2010). Increasing global temperatures and other effects of climate change could increase both the variability of these biogenic precursor emissions and/or their importance on O₃ concentrations in the United States (Demetillo et al., 2019; Gonzalez-Abraham et al., 2015; Lam et al., 2011; Lin et al., 2008, 2017; Steiner et al., 2006; Weaver et al., 2009).

The U.S.-wide reductions in anthropogenic NO_x emissions should cause O₃ in nonattainment areas to become less sensitive to variability in isoprene, and potentially more sensitive to variability in soil NO_x. The predominance of either biogenic isoprene or soil NO_x in O₃ variability is relevant for understanding where management solutions can be focused. However, to our knowledge, this policy-relevant transition in the relative importance of biogenic O₃ precursors, especially in nonattainment areas specifically, has not yet been demonstrated across the contiguous U.S. with a chemical transport model and with up-to-date emissions. While a myriad of factors can determine ambient O₃ concentrations (e.g., natural and anthropogenic emissions, long-range transport, and local meteorology), chemical transport models allow us to perform experiments where the individual drivers can be isolated from each other. For example, the direct biogenic contributions to regional air quality are often examined in experiments where the natural (or anthropogenic) emissions for a single year are turned off individually and compared to a scenario where all emissions are combined (Curci et al., 2009; Sartelet et al., 2012; Tao et al., 2003; Zare et al., 2014). However, this experimental design overlooks how interannual variability and non-additive interactions might play a role in attaining air quality targets in any given year.

In contrast to previous work which removes or adds the biogenic components to a simulation, we explore the role of the biosphere by linearizing the response of O₃ to biogenic precursors around local chemical conditions, under multiple anthropogenic NO_x emission scenarios. To do this, we test the sensitivity of simulated surface O₃ to ±50% perturbations in isoprene and soil NO_x emissions in hypothetical simulations with high and low anthropogenic NO_x emissions that are representative of declines in anthropogenic NO_x emissions in recent decades. We then combine these sensitivity experiments with predictions of biogenic emissions over a ~20-year timescale to estimate the interannual variability in summertime O₃ that is driven by these biogenic precursors in isolation from other sources of variability. We show how the importance of biogenic precursors is shifting over specific nonattainment regions across the country and discuss the relevance to policy and future development.

2. Methods

2.1. Chemical Transport Model Description

We use the GEOS-Chem v11-02 chemical transport model (<http://geos-chem.org>). GEOS-Chem simulates detailed NO_x-hydrocarbon-O₃-halogen-aerosol tropospheric chemistry, including gas–particle interactions. Our simulations are driven by assimilated meteorological data for the year 2013 provided by the NASA Global Modeling and Assimilation Office Goddard Earth Observation System. We use the GEOS-FP product which provides a nested horizontal resolution of 0.25° × 0.315° over North America and run the model with 47 vertical layers. Boundary conditions are provided by a global simulation with a 2° × 2.5° horizontal resolution. This

version of GEOS-Chem has been evaluated against observations of NO_x, O₃, and isoprene using various satellite, aircraft, and surface measurements (Fisher et al., 2016; Travis et al., 2016). Travis et al. (2016) show that the model successfully reproduces spatial variability in boundary layer O₃ measured by aircraft observations from SEAC4RS after correcting an overestimate in anthropogenic NO_x emissions, but a high bias in surface O₃ persists which is attributed to boundary layer mixing and O₃ production chemistry. Consistent with this, our baseline simulation for 2013 shows a nationwide bias in midday (12:00–04:00 p.m.) mean summertime surface O₃ concentrations compared to the CASTNET observations (NMB = 6.9 pb, $r = 0.625$). We argue that this bias does not preclude the use of GEOS-Chem in sensitivity experiments where the focus is on simulated differences in O₃ concentrations. We run our simulations from April to September. The first 2 months are used as spin-up, and we focus our results on summer months (June–July–August).

Anthropogenic emissions for the United States are provided by the National Emission Inventory (NEI-2011) for the United States as implemented by Travis et al. (2016) and by the Criteria Air Contaminants (CAC) inventory for Canada as implemented by van Donkelaar et al. (2008), respectively, with both scaled to 2013 based on economic data as described in van Donkelaar et al. (2008). Additional perturbations to the initial anthropogenic NO_x emissions are discussed below. Biogenic VOC emissions are calculated online by the Model of Emissions of Gases and Aerosols from Nature (MEGAN) v2.1 inventory (Guenther et al., 2012) as implemented by Hu et al. (2015). The biogenic emission of a compound i (E_i) is governed by the following equation:

$$E_i = E_{i,o} \times \gamma_{PAR} \times \gamma_T \times \gamma_{LAI} \times \gamma_{Age} \quad (1)$$

where $E_{i,o}$ is the base emission factor for some compound i (based on the fractional coverage of plant functional types within a grid box) under standard conditions, and deviations from these conditions are accounted for by activity factors for solar radiation (γ_{PAR}), temperature (γ_T), leaf area index (γ_{LAI}), and leaf age (γ_{Age}). Plant functional type distribution for VOC basal emission factors comes from the Community Land Model version 4 (Oleson et al., 2010), with monthly mean leaf area indices derived from Moderate Resolution Imaging Spectroradiometer (MODIS) observations (Myneni et al., 2007). Soil NO_x emissions fluxes (S_{NO_x}) are calculated online using the Hudman et al. (2012) parameterization, governed by the following equation:

$$S_{NO_x} = A'_{biome} (N_{avail}) \times f(T) \times g(\theta) \times P(l_{dry}) \quad (2)$$

where A'_{biome} is a biome-specific coefficient (which itself is a function of available soil nitrogen, N_{avail} , that accounts for the standing pool of nitrogen, deposited nitrogen from the atmosphere, and fertilizer nitrogen applications, as described in Hudman et al., 2012). This coefficient is modified by functional dependencies on temperature, $g(T)$ (where T is surface temperature), soil moisture, $g(\theta)$ (where θ is water-filled pore space), and precipitation pulsing, $P(l_{dry})$ (where l_{dry} is the dry spell length). Maps of the isoprene and soil NO_x emissions from our GEOS-Chem set-up are shown in Figure 1. Other emissions in the model include aircraft (Stettler et al., 2011), ships (Holmes et al., 2014; Vinken et al., 2011), biomass burning (GFED4.1), and lightning (Murray et al., 2012).

2.2. Sensitivity Simulations

To test the role of biogenic emissions on surface O₃, we perform two sets of sensitivity simulations: one at “high” anthropogenic NO_x emissions, and at “low” anthropogenic NO_x emissions (where anthropogenic NO_x emissions are reduced by 50%).

Nationwide anthropogenic NO_x emissions in the “high” emissions scenario are consistent with the Environmental Protection Agency's (EPA's) NEI estimate for the year 2013 (US EPA, 2021). However, many previous sources have proposed that the area and mobile anthropogenic NO_x emissions in the NEI are an overestimate (Anderson et al., 2014; McDonald, McKeen, et al., 2018; Sourì et al., 2016; Travis et al., 2016). We therefore compared our simulated NO₂ columns with observed tropospheric NO₂ columns from the Ozone Monitoring Instrument (OMI) satellite instrument over grid boxes where anthropogenic emission dominate (>75%), to estimate the time period for which these emissions are most representative. The OMI observations were accessed from <https://disc.gsfc.nasa.gov/>, filtered for pixels with a solar zenith angle smaller than 60, and cloud fractions less than 0.3. All June–July–August observations that pass these quality controls are gridded to the nested GEOS-Chem resolution

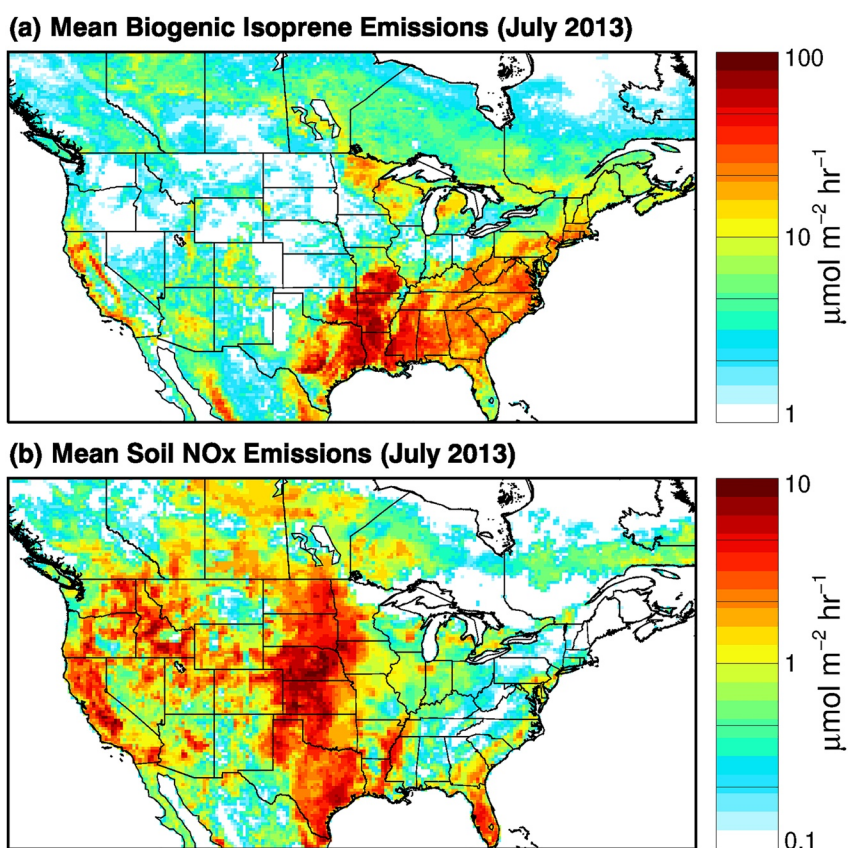


Figure 1. Mean biogenic isoprene emissions (a) and soil nitrogen oxides (NO_x) emissions (b) from GEOS-Chem for July 2013.

using the physics-based oversampling approach documented in Sun et al. (2018). By only comparing pixels where anthropogenic emissions dominate, we minimize uncertainties in this comparison introduced by lightning and soil NO_x variability, which would be affected by meteorology.

According to correlation coefficient and normalized mean bias when compared to 3 year running averages of OMI observations over urban areas, the NO₂ columns simulated by the “high anthropogenic NO_x” scenario (NEI-2013) show the best agreement with OMI observations from 2005 to 2007, while the NO₂ columns simulated by the 50% reduction in anthropogenic emissions show the best agreement with OMI observations from 2016 to 2018. Figure 2 shows this comparison of the modeled and observed tropospheric NO₂ columns (Figure S1 in Supporting Information S1 shows the full comparison of modeled columns with all 3-year running averages from 2005 to 2019). Judging from this agreement, we infer that the first set of simulations represent anthropogenic NO_x emissions most consistent with the years 2005–2007, while the second set of simulations represent anthropogenic NO_x emissions most consistent with the years 2016–2018. The model simulated O₃ in the second scenario also compares favorably with CASTNET observations for 2018 with minimal bias and a high spatial correlation (NMB = 0.1 ppb, $r = 0.86$), but we include this comparison for reference only with the caveat that we have purposely used 2013 meteorology in our simulation for the reasons described below. Our NO_x emission scenarios are therefore representative of reductions in anthropogenic NO_x emissions over the last decade and a half. This 50% decline in emissions is consistent with changes reported in other analyses over urban areas of the United States for similar time periods (Duncan et al., 2016), and with emission trends reported by the EPA (<https://www.epa.gov/air-emissions-inventories/air-pollutant-emissions-trends-data>).

For each of the high- (~2005–2007) and low- (~2016–2018) anthropogenic NO_x emission scenarios, we then perform sensitivity simulations where biogenic isoprene emissions are perturbed by +50% and –50%, respectively, and then where soil NO_x emissions are perturbed by +50% and –50%, respectively. These perturbations are accomplished by allowing GEOS-Chem to calculate the emissions online, then applying a 0.5 or 1.5 scaling

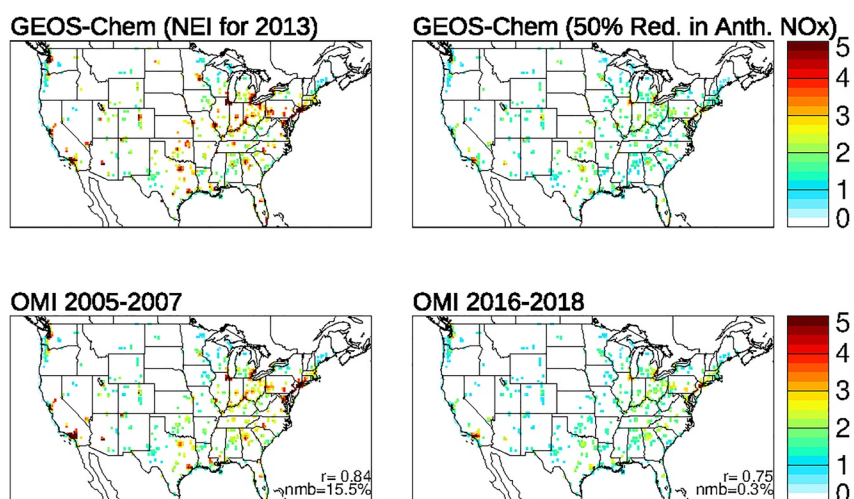


Figure 2. June–July–August mean NO_2 tropospheric column densities over model grid cells with $>75\%$ of nitrogen oxides (NO_x) emissions from anthropogenic sources. Top panels: from the GEOS-Chem simulations using NEI-2011 anthropogenic NO_x emissions scaled to 2013 (“high- NO_x ”), then a 50% reduction in anthropogenic emissions (“low- NO_x ”). Bottom panels: Ozone Monitoring Instrument (OMI) observations during the 3-year average period with the best model agreement. Pearson correlation (r) and normalized mean bias (nmb) in the model panels are shown in the bottom right corners.

factor to the calculated emissions before they are distributed into the lowest model grid box. MEGAN calculates emissions for many biogenic VOCs, and we only perturb isoprene emissions in our sensitivity experiments since our focus is on the gas-phase chemistry that impacts O_3 production (the rapid gas-phase chemistry with monoterpene and sesquiterpene emissions is not included in this model version, and this is discussed below). The $\pm 50\%$ perturbations represent reasonable and realistic regional-scale end-members given modeled long-term interannual variability in peak summertime emissions (shown below), even if the scale of interannual variability in nationwide biogenic emissions is not this large, for us to linearize the O_3 sensitivity to biogenic precursor variability around each anthropogenic emission scenario. These perturbations are also consistent with anticipated long-term changes in average emissions as a result of climate change (Gonzalez-Abraham et al., 2015).

We use the same meteorology in all our simulations to exclude minor variability in O_3 chemical sensitivity caused by meteorology, since the motivation of our study is to quantify changes in the chemical regimes resulting from emissions specifically. This provides a policy-relevant insight into the changing role of biogenic emissions, without other confounding factors. The choice of meteorological year (2013) is arbitrary. We show in Figure S2 of Supporting Information S1 that natural interannual variability alone has negligible impact on whether O_3 production is VOC- or NO_x -sensitive, so that the choice of meteorological year would not impact the results of our study. Given the sensitivity of many relevant processes to ambient temperatures, we also demonstrate in Figure S3 of Supporting Information S1 that the 2013 meteorology used here is representative of long-term mean conditions over most of the regions highlighted in our study.

We also hold anthropogenic VOC emissions constant, which artificially isolates the role of anthropogenic NO_x emissions over time. However, the NEI reports a national trend in anthropogenic VOC emissions over the United States of only -7% over the 2005–2018 period (17.753 million tons in 2005 to 16.435 million tons in 2018), so we argue this is a reasonable simplification (US EPA, 2021). It is also likely that the GEOS-Chem inventory does not accurately represent the current understanding of the importance of VOC emissions from consumer and industrial products (e.g., personal care and cleaning products) in urban areas (McDonald, De Gouw, et al., 2018; Qin et al., 2020; Coggon et al., 2021). This represents an important area of model development, but we argue that this potential omission makes our conclusions conservative since additional urban VOC emissions in our simulations could make present-day chemical conditions more NO_x -limited than currently captured by the model.

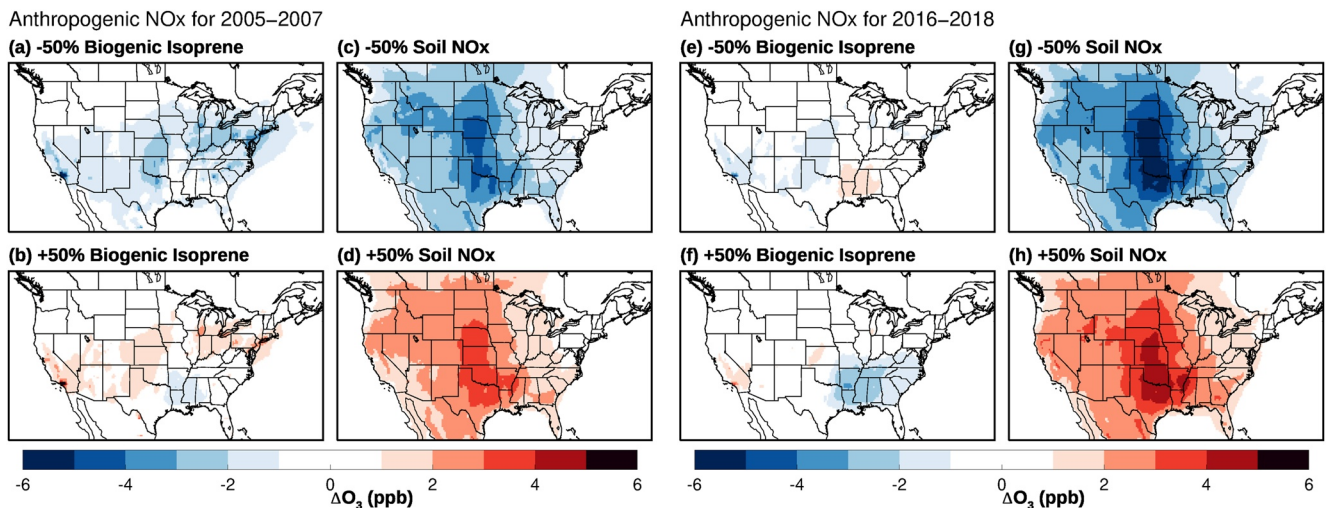


Figure 3. Sensitivity of midday (12–2 p.m. local time) mean June–July–August surface O_3 to perturbations in biogenic isoprene and soil nitrogen oxides (NOx) emissions under high anthropogenic NOx emissions (a–d) and low anthropogenic NOx emissions (e–h).

2.3. Interannual O_3 Tendencies

The impact of biogenic precursor emissions on ambient O_3 will not only depend on local-to-regional O_3 chemical sensitivity (which can be deduced from the sensitivity simulations described above), but also on the magnitude of the temporal variability in soil NOx emissions occurring year-to-year. We propose an approach where the local impact of either biogenic isoprene or soil NOx emissions on interannual variability in O_3 could be expressed as the following sum of partial derivatives:

$$\frac{dc_{O_3}}{dt} = \frac{\partial c_{O_3}}{\partial E_{isop}} \frac{\partial E_{isop}}{\partial t} + \frac{\partial c_{O_3}}{\partial E_{soilNOx}} \frac{\partial E_{soilNOx}}{\partial t} + \dots \quad (3)$$

where dc_{O_3}/dt is the interannual deviation (from some reference year) in local O_3 concentrations, $\partial c_{O_3}/\partial E$ is the local O_3 sensitivity to either isoprene or soil NOx, and $\partial E/\partial t$ is the interannual variability in their respective emission rates. Here, we call dc_{O_3}/dt the hypothetical “interannual O_3 tendency,” distinguishing it from the true interannual variability resulting from the combination of other non-linear, and often co-varying, factors, including meteorology, background O_3 , and non-additive interactions between biogenic isoprene and soil NOx emissions.

This linearization around local chemical conditions is an improvement over simulations that simply turn emissions on or off entirely. By calculating the sensitivity under different anthropogenic NOx emission scenarios, this approach provides a useful framework to demonstrate how either process impacts O_3 in isolation and under varying anthropogenic NOx emission trends. We use the sensitivity of O_3 from our perturbation simulations to estimate the $\partial c_{O_3}/\partial E_{isop}$ and $\partial c_{O_3}/\partial E_{soilNOx}$ terms for every model grid box. We determine the $\partial E_{isop}/\partial t$ and $\partial E_{soilNOx}/\partial t$ terms using output from a long-term GEOS-Chem simulation driven by structurally identical biogenic emission parameterizations (Geddes & Martin, 2017). The interannual variability in isoprene and soil NOx emissions predicted by GEOS-Chem is consistent with previous model studies and with top-down estimates inferred from satellite observations (Bauwens et al., 2016; Hudman et al., 2010; Palmer et al., 2006; Tawfik et al., 2012). Together, the estimated O_3 sensitivity results ($\partial c_{O_3}/\partial E_{isop}$ and $\partial c_{O_3}/\partial E_{soilNOx}$) and the interannual variability in emissions ($\partial E_{isop}/\partial t$ and $\partial E_{soilNOx}/\partial t$) allow us to evaluate the interannual O_3 “tendency” (dc_{O_3}/dt) resulting from either process in isolation, under the high and low anthropogenic NOx emission scenarios, respectively.

3. Results and Discussion

3.1. O_3 Sensitivity in Nonattainment Areas

Figure 3 shows the sensitivity of simulated mean midday O_3 for June–July–August resulting from a $\pm 50\%$ increase/decrease in isoprene emissions, and a $\pm 50\%$ increase/decrease in soil NOx emissions, under each of the anthropogenic NOx emission scenarios tested. Examples of zoomed detail for the southwest and NEA are

shown in Figures S4 and S5 of Supporting Information S1. In the high anthropogenic NO_x emission scenario, the perturbation to biogenic isoprene emissions had the greatest impacts on mean midday O₃ in urban areas, with some influence on regional O₃ (Figures 3a and 3b; S4 in Supporting Information S1); the effects are largest in Los Angeles, California, where O₃ concentrations change by up to ± 9 –10 ppb. In most U.S. regions, we find the isoprene response in summertime midday O₃ is fairly linear (i.e., with the positive and negative perturbations having the same magnitude of impact on O₃), except in the southeast where the effects are more non-linear. In the scenario using low anthropogenic NO_x emissions, the same $\pm 50\%$ perturbation in biogenic isoprene emissions yielded much smaller impacts as we would expect from increasingly NO_x-limited chemistry that results from a reduction in anthropogenic NO_x emissions, with O₃ unaffected in most of the urban areas (Figures 3e and 3f; S5 in Supporting Information S1). The largest response in O₃ was again observed in the Los Angeles region, but decreased to ± 4 –5 ppb, a $\sim 40\%$ – 60% reduction in sensitivity compared to the high anthropogenic emissions scenario. We note that the resolution of our chemistry simulation ($0.25^\circ \times 0.3125^\circ$) could have led us to underestimate the true VOC-limitation within cities, but the observed decrease in isoprene sensitivity suggests at least a portion of urban VOC-limitation has been captured.

Soil NO_x emission perturbations lead to changes in O₃ that are more regional. Whereas the isoprene chemical lifetime is as short as 1 hour during summer daytime (assuming OH of 0.1 ppt), the NO_x chemical lifetime under the same conditions is typically at least a few hours. In addition, NO_x can be temporarily stored in chemical reservoirs such as peroxy nitrates, such that the impacts of soil NO_x emissions on O₃ production can occur in areas farther downwind. In the scenario using high anthropogenic NO_x emissions (Figures 3c and 3d), the soil NO_x perturbation had its greatest impact on mean midday O₃ in Texas, Oklahoma, and areas of the Midwest, where O₃ changes were as high as ± 4 –5 ppb. These simulations also show a reasonably linear response in summertime midday O₃. In the low anthropogenic NO_x emission scenario (Figures 3g and 3h), the sensitivity to soil NO_x emissions both amplified and spatially expanded. The largest response in mean midday O₃ still occurred in Texas, Oklahoma, and areas of the Midwest, but increased to ± 5 –6 ppb (a $\sim 20\%$ – 25% increase in sensitivity) compared to the high anthropogenic NO_x emission scenario. We note that soil NO_x emissions make up $\sim 25\%$ of summertime NO_x emissions in the US in the “low” anthropogenic NO_x emission scenario, such that this 50% perturbation in soil NO_x only results in a $\sim 12.5\%$ change to total NO_x emissions.

To visualize the policy-relevant effects of these perturbations, we evaluated their impacts using mean daily maximum 8 hr O₃ (MDM8 O₃) averaged within the 2019 U.S. EPA-designated nonattainment areas, based on the 8 hr O₃ 2015 National Ambient Air Quality Standard of 70 ppb (shown in Figure 4a). Nonattainment areas are located predominantly in California on the west coast, the Washington-New York corridor of the NEA, and urban centers of the Great Lakes and Ohio River (GLO) valley, along with several other individual urban areas across the central, south, and southeast parts of the country. We group the nonattainment areas into regions with broadly similar biogenic precursor emissions profiles and climatic conditions: Southwest (SWE); Central (CEN); South and Southeast (SSE); GLO; and NEA.

In Figure 4, the solid lines represent the MDM8 O₃-sensitivity calculated from the high anthropogenic NO_x emission scenario, while the dashed lines represent the sensitivity calculated based on the low anthropogenic NO_x emissions. In the high anthropogenic NO_x scenario, biogenic isoprene perturbations affected O₃ in all nonattainment areas, consistent with expectations that O₃-polluted areas have been strongly VOC-limited in the past. Calculated on a population-weighted scale nationwide, O₃ in nonattainment areas varied by [+2.2, -3.1 ppb] in response to $\pm 50\%$ perturbations in biogenic isoprene emissions, compared to [+1.1, -1.3 ppb] from the $\pm 50\%$ perturbations in soil NO_x emissions. Nonattainment areas in the Southwest, GLO, and NEA regions, which together account for 78% of the U.S. population that is living in nonattainment, were the most isoprene sensitive. Therefore, despite the wider area impacts from the soil NO_x perturbations, changes in biogenic isoprene had greater impacts on O₃ where most people live, under the anthropogenic NO_x emissions consistent with ~ 2005 – 2007 levels.

In contrast, MDM8 O₃ in all nonattainment areas became more sensitive to the soil NO_x perturbations than the biogenic isoprene perturbations under the low anthropogenic NO_x emission scenario (Figure 4, dashed lines). Calculated on a population-weighted scale nationwide, O₃ in nonattainment areas varied by [+1.7, -2.0 ppb] in response to $\pm 50\%$ perturbations in soil NO_x emissions, compared to [+0.7, -1.5 ppb] from the $\pm 50\%$ perturbation in biogenic isoprene emissions. In the nonattainment areas of the NEA (Figure 4f) for example, a $\pm 50\%$ perturbation in biogenic isoprene affected MDM8 O₃ there by [+2.2, -3.3] ppb under high anthropogenic NO_x

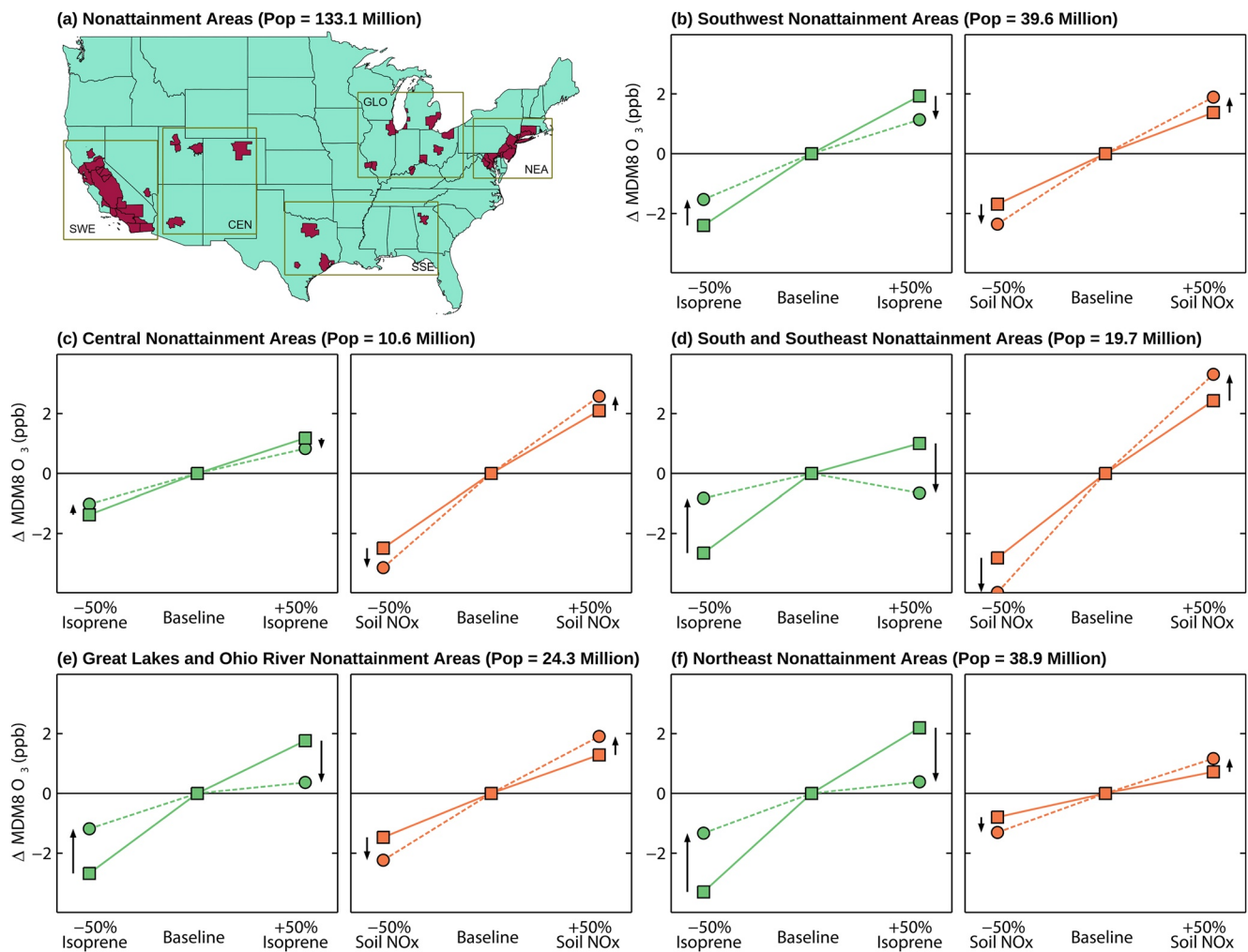


Figure 4. Sensitivity of area averaged MDM8 O₃ in U.S. nonattainment areas. (a) Map of nonattainment areas (maroon) separated by regions (rectangles). (b–f) Sensitivity of MDM8 O₃ averaged over the nonattainment areas by region, under $\pm 50\%$ perturbations to biogenic isoprene or soil nitrogen oxides (NO_x) emissions. Solid lines indicate sensitivity using an estimate of anthropogenic NO_x emissions for 2005–2007, dashed lines indicate sensitivity using an estimate anthropogenic NO_x emissions for 2016–2018 (a 50% reduction). Arrows guide the reader from the high to low NO_x anthropogenic emission scenarios.

emissions, but that sensitivity decreased by 75% to $[+0.4, -1.3]$ ppb at low anthropogenic NO_x emissions. Meanwhile, O₃ sensitivity from the $\pm 50\%$ perturbation in soil NO_x emissions in the nonattainment areas of this region increased from $[+0.7, -0.8]$ to $[+1.2, -1.3]$ under high and low anthropogenic NO_x emissions, respectively.

In SSE nonattainment areas, the sensitivity of MDM8 O₃ to the $\pm 50\%$ biogenic isoprene perturbations switched signs between the two anthropogenic emission scenarios, varying from $[+1.0, -2.7]$ ppb under high anthropogenic NO_x emissions to $[-0.7, -0.8]$ ppb at low anthropogenic NO_x emissions. Meanwhile, O₃ sensitivity from the $\pm 50\%$ perturbation in soil NO_x emissions in the nonattainment areas of this region increased from $[+2.4, -2.8]$ to $[+3.3, -4.0]$ under high and low anthropogenic NO_x emissions, respectively. Therefore, despite the conventional expectation of VOC-limitation in urban areas, biogenic soil NO_x has greater impacts on O₃ where most people live, under anthropogenic NO_x emissions consistent with ~ 2018 levels.

We note that the 50% reduction in anthropogenic NO_x emissions alone results in a 6 ppb decrease in population weighted mean MDM8 O₃ across the nonattainment areas nationwide in our simulations (compared to ~ 2 ppb decreases in population weighted mean MDM8 O₃ from the 50% reduction in soil NO_x or biogenic isoprene), demonstrating the principal role that steady declines in anthropogenic NO_x emissions has had overall on O₃ air quality in the United States.

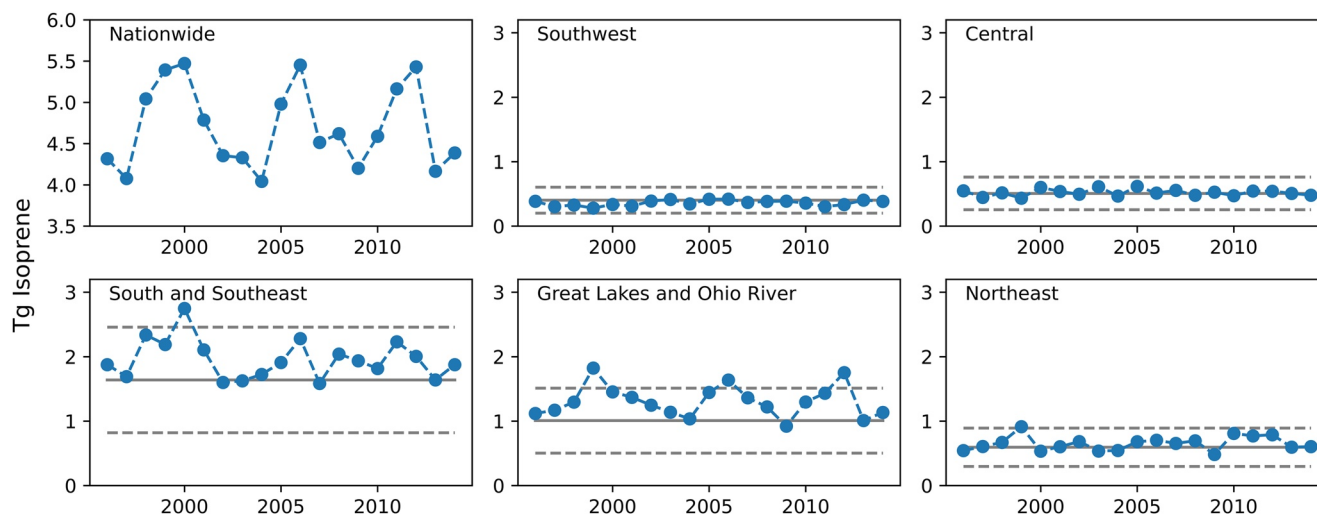


Figure 5. Interannual variability in July biogenic isoprene emissions nationwide, and for each of the individual regions highlighted in rectangles in Figure 4. The gray solid line denotes the year 2013 emissions (which are our baseline), and the dashed lines denote the $\pm 50\%$ envelope used in the perturbation experiments.

3.2. Interannual Tendencies of O_3 in Nonattainment Areas

As described in the methods, the impact of biogenic precursor emissions on observed O_3 will not only depend on local-to-regional O_3 chemical sensitivity (which can be deduced from the experiments described above), but also on the magnitude of the temporal variability in soil NO_x emissions occurring year-to-year. Here, we use the interannual variability in soil NO_x and biogenic isoprene emissions calculated from a long-term GEOS-Chem simulation (Geddes & Martin, 2017) whose biogenic emissions are structurally identical to the nested model setup used here. July isoprene emissions over the United States in this model are on average 4.7 ± 0.5 , ranging from as low as 4.0 Tg (in 2004) to as high as 5.5 Tg (in 2000), a difference of almost 40% relative to the minimum (or $\pm 15\%$ relative to the long-term mean). The scale of interannual variability is consistent with the scale of interannual variability in previous MEGAN model studies (Tawfik et al., 2012) and inferred from top-down estimates based on satellite-observed formaldehyde (an atmospheric oxidation product of isoprene) (Bauwens et al., 2016; Palmer et al., 2006). July soil NO_x emissions in this model are on average $0.15 \text{ Tg} \pm 0.02$, and range from as low as 0.11 Tg (in 1997) to 0.20 Tg (in 2012), a difference of 82% relative to the minimum (or $\pm \sim 30\%$ relative to the long-term national mean). These results are consistent with estimates of variability derived from satellite observations (Hudman et al., 2010).

Figures 5 and 6 show the timeseries of nationwide and regional biogenic isoprene and soil NO_x emissions for the month of July as predicted by the model. We find that our choice of $\pm 50\%$ relative to our baseline 2013 conditions allows us to reasonably capture natural variability that does occur the regional scale. It is of course difficult to come up with a perturbation experiment that will be exactly representative of contemporary interannual variability in the emissions everywhere: in some cases, only the response in one direction (e.g., +50%, or -50%) is important; in other cases, this envelope can underestimate of the local variability (e.g., soil NO_x variability in the SSE, and in the GLO regions). We are most interested in interpolating the response of O_3 between these two end members to estimate the O_3 response to natural interannual variability (Section 3.2), so the exact choice is less important than being able to fairly interpolate between the end members.

Using Equation 3 to linearize O_3 sensitivity around local chemical conditions, we calculate the interannual July mean O_3 tendencies driven by either soil NO_x or biogenic isoprene emission variability from year to year under the “high” (2005–2007) and “low” (2016–2018) anthropogenic NO_x emissions. To visualize the relative importance of biogenic isoprene and soil NO_x emissions in the July mean interannual O_3 tendencies, we define a normalized difference index as:

$$\frac{\sigma_{\text{soil NO}_x} - \sigma_{\text{isoprene}}}{\sigma_{\text{soil NO}_x} + \sigma_{\text{isoprene}}}$$

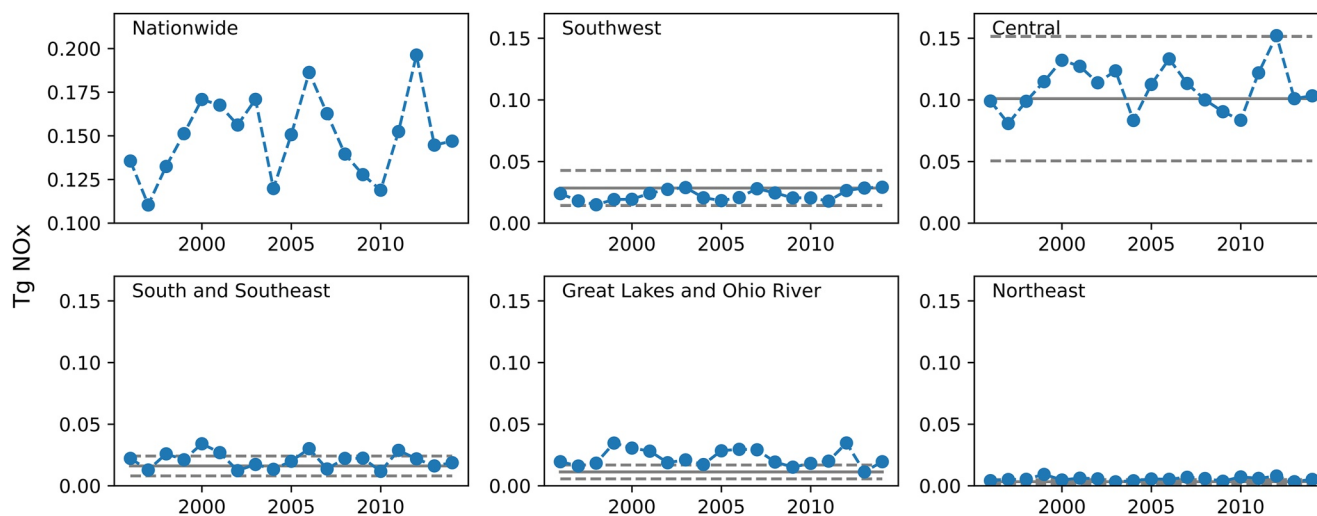


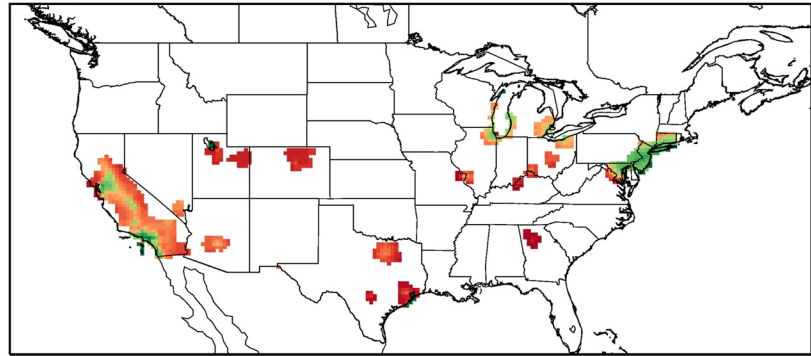
Figure 6. Interannual variability in July soil nitrogen oxides (NOx) emissions nationwide, and for each of the individual regions highlighted in rectangles in Figure 4. The gray solid line denotes the year 2013 emissions (which are our baseline), and the dashed lines denote the $\pm 50\%$ envelope used in the perturbation experiments.

where $\sigma_{\text{soil NOx}}$ is the standard deviation in the interannual tendencies of O_3 from soil NOx variability, and σ_{isoprene} is the standard deviation in the interannual tendencies of O_3 from soil NOx variability. Given this definition, a value of +1 indicates that soil NOx emissions dominate the standard deviation of the interannual tendencies while a value of -1 indicates that biogenic isoprene emissions dominate. Values closer to 0 mean that each are equally important in the interannual O_3 tendencies.

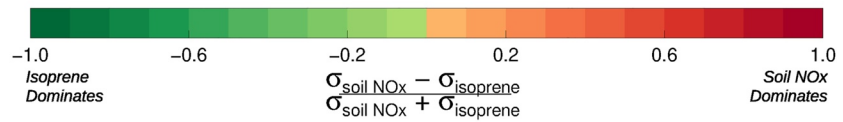
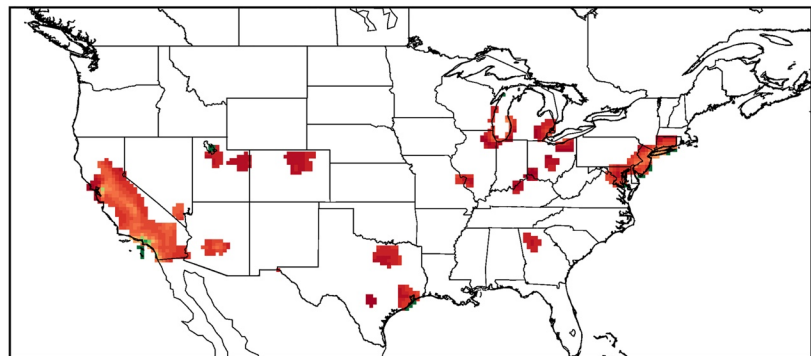
Figure 7 maps the normalized difference index under each hypothetical NOx emission scenario. We find that biogenic isoprene drives the interannual tendencies in several of the most populated nonattainment areas under the high anthropogenic NOx emission scenario, but that soil NOx drives the interannual tendencies in all nonattainment areas in the low anthropogenic NOx emission scenario, with some exceptions in California. In Figures 7c–7e, we present the distribution of interannual O_3 tendencies from three specific nonattainment locations: Los Angeles (Southwest), Dallas (Central), and Philadelphia (NEA). In all three examples, the role of biogenic isoprene emissions on July-mean interannual O_3 tendencies diminishes with the reduction in anthropogenic NOx emissions, while the contribution from soil NOx emissions increases. In Los Angeles, the dominance of variability in biogenic isoprene over soil NOx persists in both scenarios. In Dallas, which is representative of the nonattainment areas of the SSE region, soil NOx is more important than biogenic isoprene in the interannual O_3 tendencies under both scenarios. Finally, in Philadelphia, which is representative of nonattainment areas in the NEA region, the importance of biogenic isoprene versus soil NOx on interannual O_3 tendencies switches such that soil NOx now dominates low anthropogenic NOx scenario. This shift from isoprene-dominant interannual O_3 tendencies to soil-NOx dominated interannual O_3 tendencies throughout the nonattainment areas of the Great Lakes, Ohio River, and NEA regions (Figure 5a) is noteworthy. In these parts of the country where chemistry had traditionally been VOC limited and sensitive to interannual variability in biogenic isoprene, we expect interannual variability in soil NOx to become more important.

While idealized, these “ O_3 tendencies” demonstrate how, given the observed reductions in anthropogenic NOx emissions, we can expect the importance of biogenic isoprene on O_3 interannual variability to decrease in many nonattainment regions across the country if anthropogenic NOx reductions continue. At the same time, the O_3 tendencies suggest O_3 interannual variability in nonattainment areas will be increasingly sensitive to soil NOx emissions variability. We note that the reduction in anthropogenic NOx in our simulations had a larger overall impact on O_3 than the interannual variability inferred from these biogenic precursors, demonstrating the more important role that steady declines in anthropogenic NOx have had on O_3 over the long term. But this analysis shows that these precursors are, and will continue to be, important to interannual variability in ways that are changing in response to the anthropogenic emission declines. In the model grid box over Philadelphia, for example, soil NOx emissions comprise $\sim 11\%$ of the local NOx emissions in July (using anthropogenic NOx emissions consistent with 2016–2018). The variability in total NOx emissions here due to the interannual variability in soil

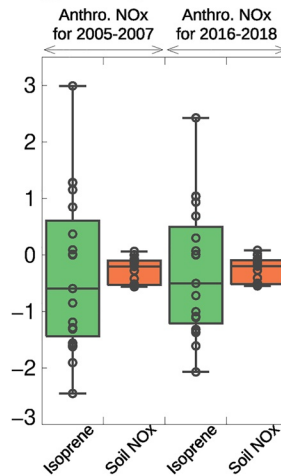
(a) Anthropogenic NOx for 2005-2007



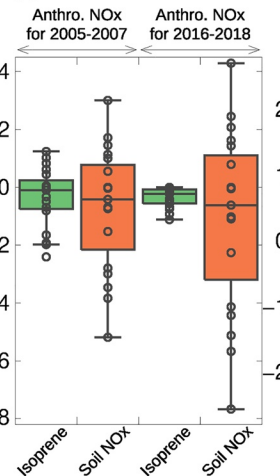
(b) Anthropogenic NOx for 2016-2018



(c) Los Angeles



(d) Dallas



(e) Philadelphia

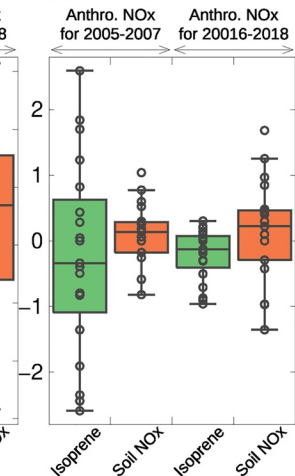


Figure 7. (a and b) Increasing importance of soil nitrogen oxides (NOx) processes on O_3 interannual tendencies across nonattainment areas of the United States using a normalized difference index (where σ is the standard deviation in interannual tendencies); and (c–e) the distribution of interannual July mean O_3 tendencies due to interannual variability in biogenic isoprene and soil NOx emissions for three specific nonattainment grid boxes, assuming O_3 sensitivity using high (~2005–2007) anthropogenic NOx emissions (green bars) versus O_3 sensitivity using low (~2016–2018) anthropogenic NOx emissions (orange bars).

NO_x emissions is therefore quite small locally, with interquartile range of [1%, 4%] relative to 2013; still, these perturbations cause a change in mean July O₃ relative to 2013 with an interquartile range of [0.29 ppb, 0.48 ppb], Figure 5f).

Our model experiments are also relevant to understanding an “O₃-climate penalty” (where the impact of emission controls on O₃ concentrations can be offset due to increasing environmental temperatures). The magnitude and mechanisms behind a climate penalty are sensitive to the local chemical environment, and thus to changes in anthropogenic emissions over time (Bloomer et al., 2009; Rasmussen et al., 2013). The results of our perturbation studies shed light on the specific role individual biogenic emissions can play in the O₃-climate penalty over non-attainment areas in the U.S., and how these roles evolve with changes in anthropogenic emissions. In the Philadelphia region for example, the interannual O₃ tendencies due to isoprene correlate strongly with temperature in both scenarios (from $r = 0.97$ to $r = 0.91$), but the interannual tendencies from soil NO_x (which are also strongly correlated with temperature at that location, $r = 0.76$) dominate when anthropogenic NO_x emissions are reduced, so that the absolute role of an isoprene-temperature correlation is less important.

Additional factors will contribute to the true O₃ variability depending on location, including meteorology, and chemistry; but we interpret these results to mean that soil NO_x emissions will become an increasingly important control on high-O₃ days in many U.S. nonattainment areas. Environmental and land management factors play a role in the timing and magnitude of soil NO_x emissions (Pilegaard, 2013; Stehfest & Bouwman, 2006; Steinkamp & Lawrence, 2011) in ways that differ from our understanding of the timing and magnitude of isoprene emissions. The dependence of soil NO_x emissions on nitrogen availability (including fertilizer application), soil water content (including precipitation timing and the preceding dry spell length), and soil pH could introduce new sub-seasonal and/or episodic patterns in high O₃ that may not have historically been considered in nonattainment areas where chemistry has for a long time been VOC-limited. Agricultural land that is actively managed by humans contributes a disproportionate fraction of soil NO_x emitted in the United States (Almaraz et al., 2018; Davidson et al., 1998). In our GEOS-Chem simulation, almost 40% of total soil NO_x emissions in the United States occur over cropland or mixed cropland/natural vegetation mosaic, with 30% of total soil NO_x emissions associated with fertilizer application. Almaraz et al. (2018) outline several approaches that could potentially be used to reduce soil NO_x emissions from fertilized croplands (e.g., slow-release fertilizers, precision agriculture, and cover crops), and our results here provide further motivation to explore these strategies.

We note that our results also imply the importance of constraining our estimates of biogenic isoprene and soil NO_x emissions. Several methods exist for predicting both biogenic isoprene and soil NO_x emissions, and some aspects of our inferred O₃ sensitivity may be specific to the parameterizations used (Carlton & Baker, 2011; Hogrefe et al., 2011; Kim et al., 2017; Rasool et al., 2019). While uncertainties in biogenic isoprene emissions may be of decreasing importance as they relate to interannual O₃ nonattainment in the United States, it is becoming increasingly important to reduce uncertainties in soil NO_x variability. Further constraints on the variability of soil NO_x emissions, such as those informed by satellite-based observations (Hudman et al., 2010; Vinken et al., 2014), will be valuable. Sha et al. (2021) and Wang et al. (2021) recently showed that a revised version of the soil NO_x parameterization scheme (mainly, an updated temperature response) tends to agree better with satellite observations of NO₂ over rural areas of the United States. More detail may be gathered from next-generation satellites that have improved spatial and temporal resolution.

Advancing our mechanistic understanding of NO_x uptake by vegetation canopies (Delaria & Cohen, 2019) and within-canopy nitrogen chemistry (Min et al., 2014), improving our understanding of microbial controls over soil NO_x emissions from natural landscapes including forests (Mushinski et al., 2019), and quantifying soil NO_x emissions from bare or vegetated urban ground where atmospheric nitrogen inputs are particularly high (Decina et al., 2017; Rao et al., 2014) may be critical. Continued progress in representing isoprene chemistry (Bates & Jacob, 2019), as well as including other biogenic VOC gas-phase chemistry mechanisms (Fisher et al., 2016; Porter et al., 2017), is also important as O₃ sensitivity could be influenced by the choice of chemical mechanism (Dunker et al., 2016; Squier et al., 2014), though our results are spatially consistent with an earlier GEOS-Chem sensitivity experiment (Mao et al., 2013). Finally, this work highlights the importance of adopting accurate, up-to-date, and high resolution land cover maps in chemical transport simulations, with timely information about agricultural land management (Rasool et al., 2019). Despite the aspects of model development described above that would contribute to efficient O₃ policies going forward, our results provide a reasonable range of impact and highlight the emerging influence of soil NO_x emissions on O₃ in U.S. nonattainment areas.

4. Conclusions

U.S. urban areas are shifting to increasingly NO_x-limited O₃ production, and we will therefore encounter greater O₃ sensitivity to interannual variability in soil NO_x. In this study, we quantify this changing sensitivity by comparing the importance of biogenic precursors in a chemical transport model at high and low anthropogenic NO_x emissions representative of declines observed in recent decades. Biogenic isoprene emissions are usually considered beyond regulatory control, so the finding that they are becoming less influential on surface O₃ variability in nonattainment areas is useful for air quality management. On the other hand, there is growing importance of soil NO_x emissions on O₃ chemistry in nonattainment areas that is as large as isoprene effects 15 years ago, both by area-averaged and population-weighted metrics.

Understanding how soil NO_x emissions will increasingly contribute to high O₃ episodes, in addition to season-long interannual variability, may be required in designing effective strategies for reducing the number of bad O₃ air quality days in the future. While our experiment focuses on the traditional summer O₃ season, O₃ soil NO_x-sensitivity could have implications for the seasonal O₃ cycle and the timing of O₃ extremes within seasons, potentially varying with agricultural phenology, land management, and precipitation or soil wetness, which will be the subject of future work.

Data Availability Statement

The GEOS-Chem model output used in this study is freely accessible via the OpenBU Research Data repository (<https://open.bu.edu/handle/2144/44806>). This repository provides open access and long-term digital preservation of all data.

Acknowledgments

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