



OPEN ACCESS

EDITED BY

Günther Thiele,
Freie Universität Berlin, Germany

REVIEWED BY

Guo-Hua Zhong,
Chinese Academy of Sciences (CAS),
China
Yong Du,
Merck, United States

*CORRESPONDENCE

Luliia Koemets,
✉ koemets.j@gmail.com
Leonid Dubrovinsky,
✉ leonid.dubrovinsky@uni-bayreuth.de

RECEIVED 13 July 2023

ACCEPTED 19 September 2023

PUBLISHED 06 October 2023

CITATION

Koemets I, Wang B, Koemets E, Ishii T,
Liu Z, McCammon C, Chanyshv A,
Katsura T, Hanfland M, Chumakov A and
Dubrovinsky L (2023), Crystal chemistry
and compressibility of
 $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ and $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$
silicate perovskites at pressures up
to 95 GPa.
Front. Chem. 11:1258389.
doi: 10.3389/fchem.2023.1258389

COPYRIGHT

© 2023 Koemets, Wang, Koemets, Ishii,
Liu, McCammon, Chanyshv, Katsura,
Hanfland, Chumakov and Dubrovinsky.
This is an open-access article distributed
under the terms of the [Creative
Commons Attribution License \(CC BY\)](#).
The use, distribution or reproduction in
other forums is permitted, provided the
original author(s) and the copyright
owner(s) are credited and that the original
publication in this journal is cited, in
accordance with accepted academic
practice. No use, distribution or
reproduction is permitted which does not
comply with these terms.

Crystal chemistry and compressibility of $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ and $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ silicate perovskites at pressures up to 95 GPa

Luliia Koemets^{1*}, Biao Wang², Egor Koemets², Takayuki Ishii³,
Zhaodong Liu⁴, Catherine McCammon¹, Artem Chanyshv¹,
Tomo Katsura¹, Michael Hanfland⁵, Alexander Chumakov⁵ and
Leonid Dubrovinsky^{1*}

¹Bayerisches Geo Institute (BGI), Universität Bayreuth, Bayreuth, Germany, ²Department of Earth Sciences, University of Oxford, Oxford, United Kingdom, ³Institute for Planetary Materials, Okayama University, Misasa, Japan, ⁴State Key Laboratory of Superhard Materials, Jilin University, Changchun, China, ⁵European Synchrotron Radiation Facility (ESRF), Grenoble, France

Silicate perovskite, with the mineral name bridgmanite, is the most abundant mineral in the Earth's lower mantle. We investigated crystal structures and equations of state of two perovskite-type Fe^{3+} -rich phases, $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ and $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$, at high pressures, employing single-crystal X-ray diffraction and synchrotron Mössbauer spectroscopy. We solved their crystal structures at high pressures and found that the $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ phase adopts a novel monoclinic double-perovskite structure with the space group of $P21/n$ at pressures above 12 GPa, whereas the $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ phase adopts an orthorhombic perovskite structure with the space group of $Pnma$ at pressures above 8 GPa. The pressure induces an iron spin transition for Fe^{3+} in a $(\text{Fe}_{0.7}\text{Mg}_{0.3})\text{O}_6$ octahedral site of the $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ phase at pressures higher than 40 GPa. No iron spin transition was observed for the $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ phase as all Fe^{3+} ions are located in bicapped prism sites, which have larger volumes than an octahedral site of $(\text{Al}_{0.5}\text{Si}_{0.5})\text{O}_6$.

KEYWORDS

bridgmanite, silicate perovskite, double perovskite, spin transition, single-crystal X-ray diffraction, synchrotron Mössbauer spectroscopy, high pressure

1 Introduction

The most abundant mineral on the Earth, magnesium silicate perovskite (bridgmanite), crystallizes in an orthorhombic GdFeO_3 -type perovskite structure and consists of large distorted “bicapped prism” sites (pA-sites) in the voids of the three-dimensional net of corner-sharing octahedra (oB-sites) (Figure 1C). Compressibility (Fiquet et al., 2000; Tsuchiya et al., 2004; Vanpeteghem et al., 2006) and Brillouin spectroscopy (Sinogeikin et al., 2004) studies on the MgSiO_3 bridgmanite end-member reported isothermal bulk modulus values ranging between 259 and 268 GPa. The effect of Al and Fe content on bridgmanite compressibility remains unclear due to limited information on the substitution mechanisms, iron oxidation state (Mao et al., 1991; Kubo et al., 2000; Andraut et al.,

2001; Nishiyama et al., 2007), and cation distribution in the samples (Saikia et al., 2009). The Fe^{3+} content may be significant in bridgmanite even at low oxygen fugacity ($f(\text{O}_2)$) (Frost et al., 2004), and Fe^{3+} may cause significant changes in the elastic properties of the material (Saikia et al., 2009; Ballaran et al., 2012). Moreover, the presence of oxygen vacancies strongly decreases the bulk modulus (Ismailova et al., 2016); therefore, substitution mechanisms occurring during sample synthesis should be considered (Mao et al., 2015; Shukla and Wentzcovitch, 2016).

Fe- and Al-rich sample syntheses for further diamond anvil cell (DAC) experiments are challenging because it is difficult to obtain homogeneous compositions with sufficiently large crystals (up to micrometers) of a quality suitable for single-crystal X-ray diffraction (SC-XRD). Therefore, previous X-ray diffraction experiments at high pressures were usually limited by low Fe and Al contents (Ballaran et al., 2012; Glazyrin et al., 2014) in the sample or with the use of powder diffraction, leading to complicated data interpretation for the studies of silicate crystal chemistry at high pressure (Liu et al., 2018; Zhu et al., 2020).

In the present study, we synthesized high-quality crystals of high-pressure silicates with high Fe and Al contents, employing a multi-anvil apparatus. The samples were further loaded in diamond anvil cells for *in situ* SC-XRD experiments up to 60 GPa and the Mössbauer spectroscopy study at pressures up to 95 GPa. We unambiguously identified the structure of high-pressure Fe-bearing Al-free silicate $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ as double perovskites with two octahedral sites, one (oB) occupied by silicon and another (oB') by ferric iron and magnesium. We were able to observe the volume collapse of the Fe^{3+} -bearing oB'-site and changes in Mössbauer parameters at pressures above 40 GPa, which were previously associated with the spin transition. We were also able to derive the "FeAlO₃" end-member bulk modulus. Our results show that compositional variations in bridgmanite have an impact on the structure and crystal chemistry and lead to the appearance of a more complex and unusual phase, silicate double perovskites.

2 Materials and methods

2.1 Sample synthesis and characterization

$\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ and $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ single crystals were synthesized using the Kawai-type multi-anvil press with the Osugi-type module (Ishii et al., 2019) at Bayerisches Geoinstitut, IRIS-15 (Ishii et al., 2016). A detailed description of the sample synthesis procedure can be found in Liu et al. (2019). The chemical composition of the recovered samples was determined using a JEOL JXA-8200 Electron Probe Microanalyzer (EPMA). The oxidation state of iron was determined by Mössbauer spectroscopy. Within the detection limits of the measurements, all the iron in $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ is represented as Fe^{3+} . Approximately 16(4)% of iron in $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ is represented as Fe^{2+} (Supplementary Figure S1).

2.2 High-pressure experiments

BX90-type (Kantor et al., 2013) diamond anvil cells with diamond culet sizes ranging from 120 to 250 μm were used for conducting high-pressure experiments. To create a sample chamber

between diamonds, rhenium gaskets were pre-indented to a thickness of $30 \pm 5 \mu\text{m}$. Subsequently, a laser was used to drill a hole in the center of the indented area, creating a sample chamber with a diameter of 50–110 μm depending on diamond culet sizes. Pre-selected single crystals were loaded into the center of the sample chamber together with a ruby sphere for pressure determination (Mao et al., 1986) at low pressures ($p < 6 \text{ GPa}$). Neon gas was loaded (Kurnosov et al., 2008) around the samples to serve as pressure-transmitting mediums, minimizing the degree of deviatoric stress. Additionally, it was also used for pressure determination at high pressures ($p > 6 \text{ GPa}$; Fei et al., 2007).

2.3 Single-crystal X-ray diffraction

SC-XRD patterns were collected at the ID15B beamline at the European Synchrotron Radiation Facility (ESRF). An X-ray beam with the energy of 30 keV ($\lambda = 0.4133 \text{ \AA}$) was used, and diffraction data were collected using a MAR555 flat-panel detector. At each pressure point, SC-XRD data collection was performed in the omega range of $\pm 38^\circ$ or $\pm 32^\circ$, depending on the DAC opening angle, with a 0.5° step and exposure time of 1 s for each step. The integration of the reflection intensities and absorption corrections was performed using CrysAlis^{pro} (Agilent, 2014). The structure solution and refinement were performed in the isotropic approximation using Jana2006 (Petříček et al., 2014) with Superflip (Palatinus and Chapuis, 2007) and SHELXT (Sheldrick, 2015).

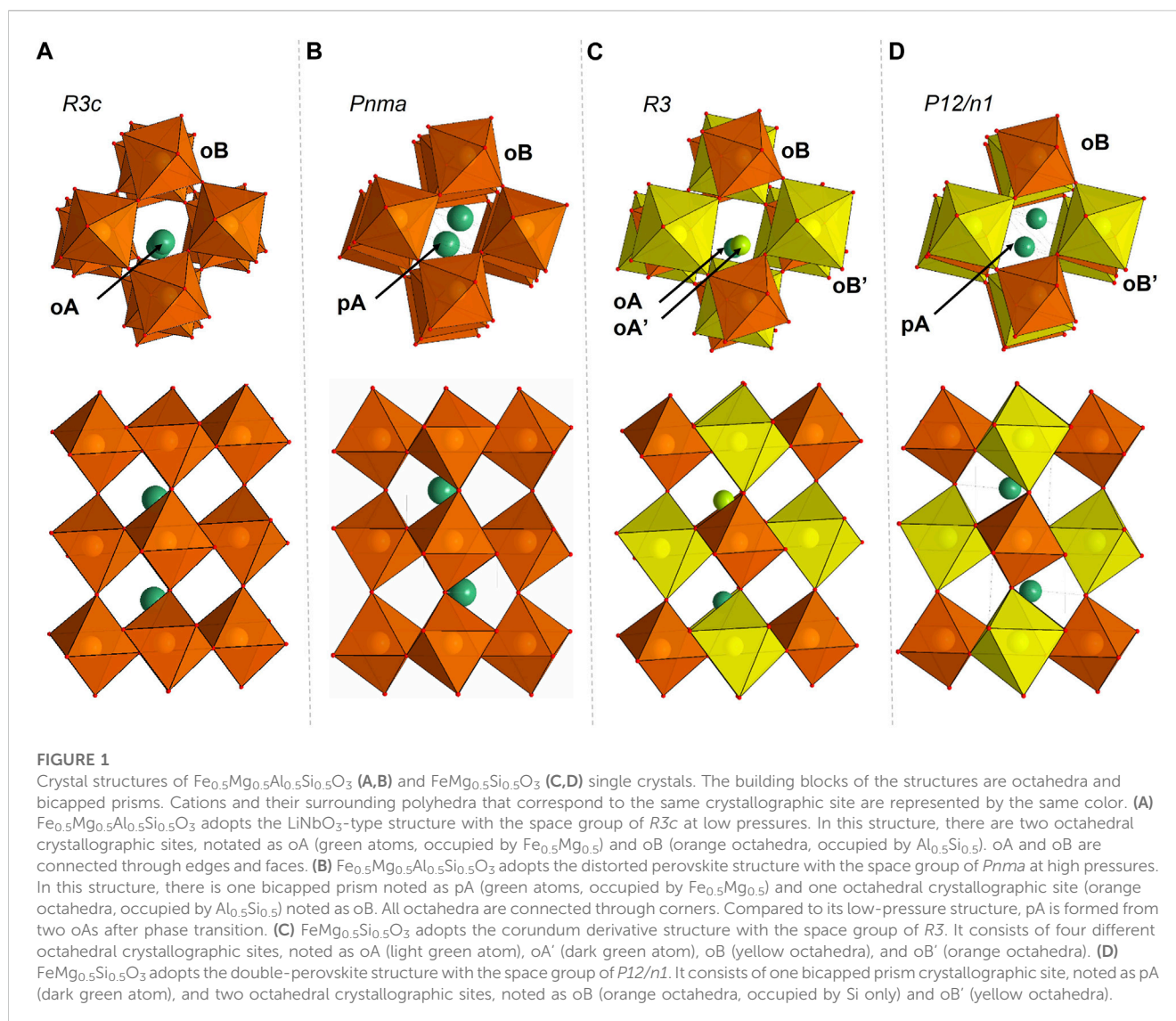
2.4 Synchrotron Mössbauer spectroscopy

Energy-domain synchrotron Mössbauer spectroscopy measurements were carried out at the nuclear resonance beamline ID18 at ESRF (Rüffer and Chumakov, 1996), using the synchrotron Mössbauer source (Potapkin et al., 2012). The spot size of the focused beam was approximately $15 \mu\text{m}^2 \times 15 \mu\text{m}^2$. Due to the usage of ^{57}Fe in the starting material during sample synthesis, spectral acquisition times were less than 1 h. Therefore, we do not expect an appearance of spectral features associated with the signal obtained from Fe contained in the Be window and lenses, as stated in the previous studies.

3 Results and discussion

3.1 Crystal structures of $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$

At ambient conditions, $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ adopts a LiNbO_3 -type structure (space group $R3c$; Figure 1). The lattice parameters are determined to be $a = b = 4.8790(1) \text{ \AA}$ and $c = 12.9112(1) \text{ \AA}$. This structure consists of two types of octahedra, oA and oB, forming corundum-like layers stacked along the crystallographic c axis (Ishii et al., 2017). Octahedral oA-sites are occupied by Fe and Mg (with an atomic ratio of 1:1), and oB-sites are occupied by Al and Si (with an atomic ratio of 1:1). On compression, $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ is stable in the LiNbO_3 -type structure up to $8 \pm 2 \text{ GPa}$ (Figure 4), above which it transforms into a distorted perovskite structure (space group $Pnma$; Figures 1B, 4). As expected from the difference in ionic radii in the



octahedral coordination (Shannon, 1976), the volume of oA octahedra is larger than that of oB (Figure 5). The phase transitions occur through the tilt of oB octahedra and a shift of oA-site cations to form eight-fold prismatic sites (pA in Figures 1A, C).

3.2 Crystal structures of $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$

Compared to $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$, $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ exhibits a more complex structure, as supported by the systematic absence analysis (Supplementary Figure S2). At ambient conditions, it adopts a corundum derivative structure (space group $R3$; Figure 1C), with the lattice parameters determined to be $a = b = 4.9406(7)$ Å and $c = 13.319(2)$ Å. Two octahedra, labeled as oA and oA', are occupied by Fe and Mg, with site occupancies of $(\text{Fe}_{0.6}\text{Mg}_{0.4})$ and $(\text{Fe}_{0.7}\text{Mg}_{0.3})$, respectively, and are located in every second layer along the c -direction. Another two octahedra, oB and oB', are occupied by Si and $(\text{Mg}_{0.7}\text{Fe}_{0.3})$. The difference in oB- and oB'-site occupancies causes the difference in octahedra volumes at ambient conditions: oB has a

volume of 8 \AA^3 , while oB' has a volume of 10.4 \AA^3 . It is worth noting that oB and oB' are located in layers along c -directions and do not mix with the layers of oA and oA' octahedra. It is a new structure, which has not been observed previously for corundum derivatives. Upon compression, $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ is stable in the corundum derivative structure up to 12 ± 2 GPa, above which it transforms into a double-perovskite structure (space group $P12/n1$; Figures 1D, 4). This structure has some unique features that were not observed in other compositions previously. First, we observed the ordering of cations located on octahedral sites (oB and oB'), which leads to a symmetry decrease from orthorhombic perovskites to monoclinic double perovskites (Figure 1D). As a result, two different oB- and oB'-sites remain distinguishable after phase transition at 12 GPa (Figure 5). Second, we see that a larger pA-site is occupied by Mg and Fe based on single-crystal X-ray diffraction data refinement. The average pA-site Fe^{3+} occupancy is approximately 0.62(3) and is kept for all measured pressure points (Supplementary Figure S3). Another oB-site is occupied only by Si, while the oB'-site contains Fe and sufficient amounts of Mg (~30%). The derived cation distribution is not an artifact because an attempt to refine structures within the oB-site occupied by both Si and Fe,

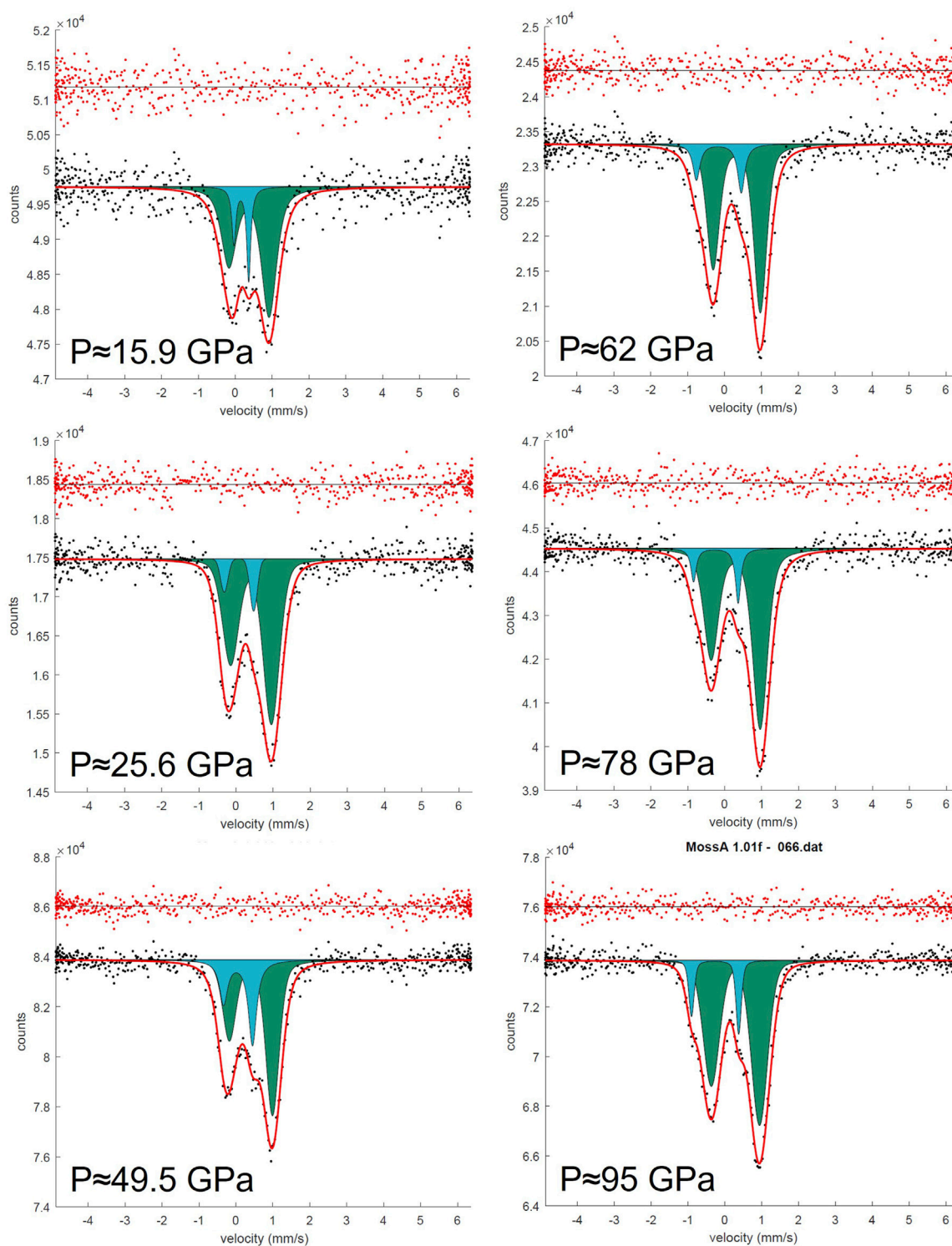
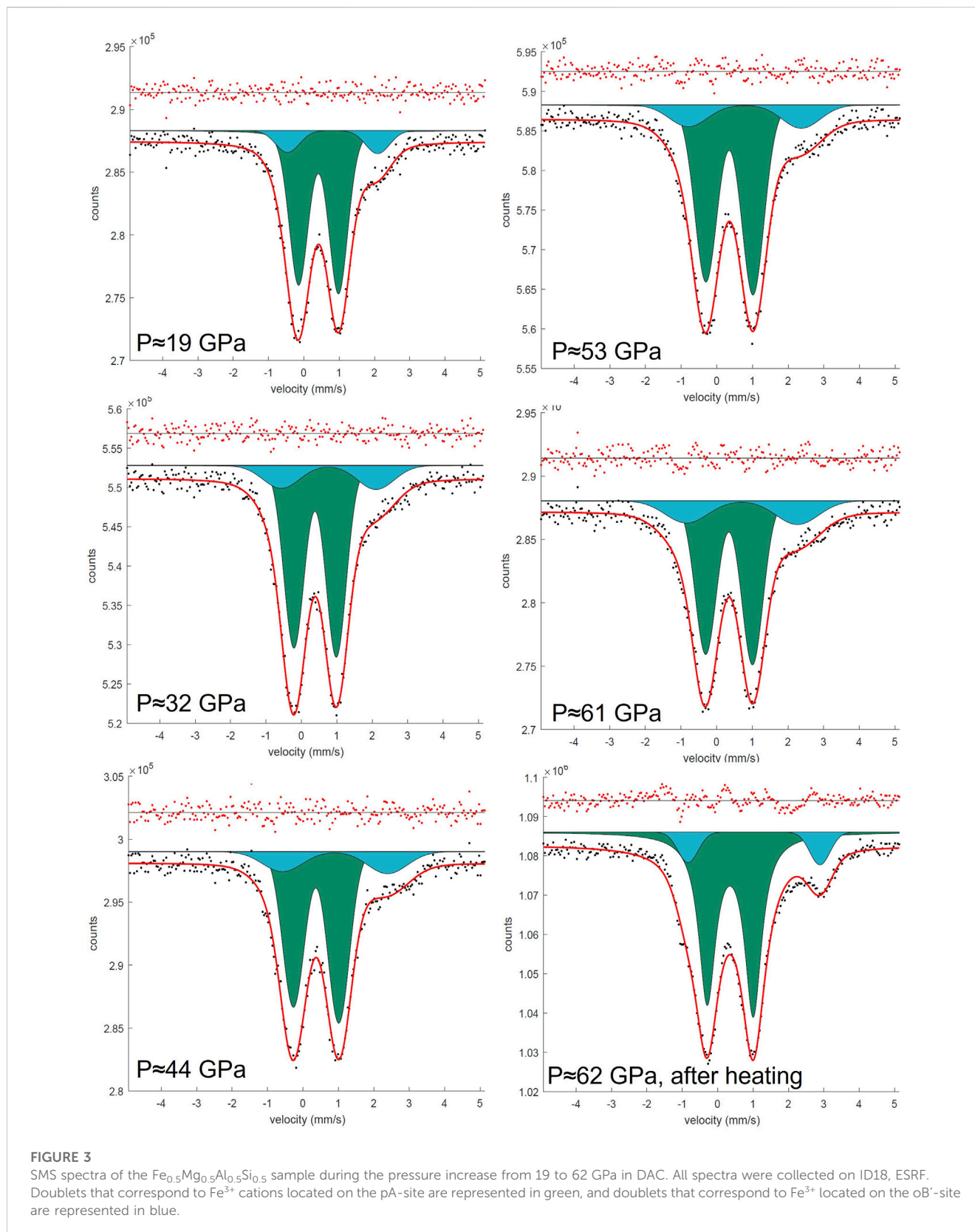


FIGURE 2

SMS spectra of the $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ sample during the pressure increase from 15.9 to 95 GPa in DAC. All spectra were collected on ID18, ESRF. Doublets that correspond to Fe^{3+} cations located on the pA-site are represented in green, and doublets that correspond to Fe^{3+} located on the oB'-site are represented in blue.

or oB' without the Mg-worth refinement quality (R_{all} increase; see [Supplementary Figure S4](#) for more information). Considering all the available data, we conclude that, in all previously investigated

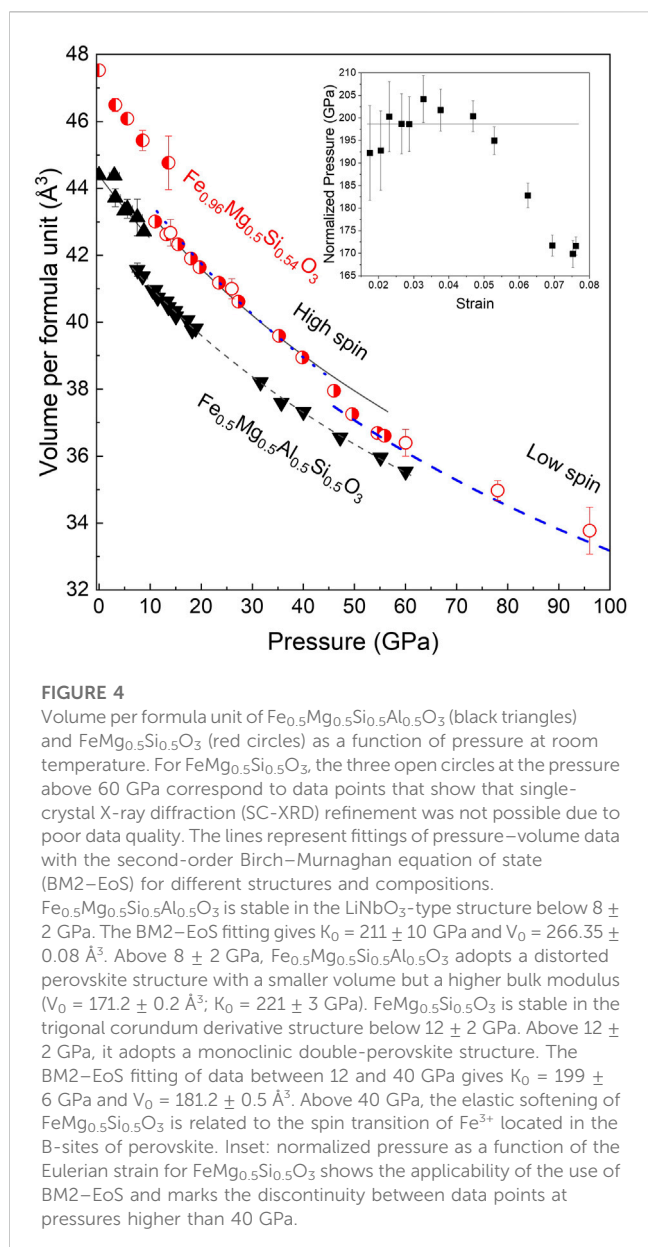
compositions, the pA-site is occupied by $\text{Fe}_{0.6}\text{Mg}_{0.4}$ and the oB'-site is occupied by $\text{Fe}_{0.7}\text{Mg}_{0.3}$. More information on structure refinement can be found in [Supplementary Tables S1–S3](#).



3.3 Iron: oxidation and spin states

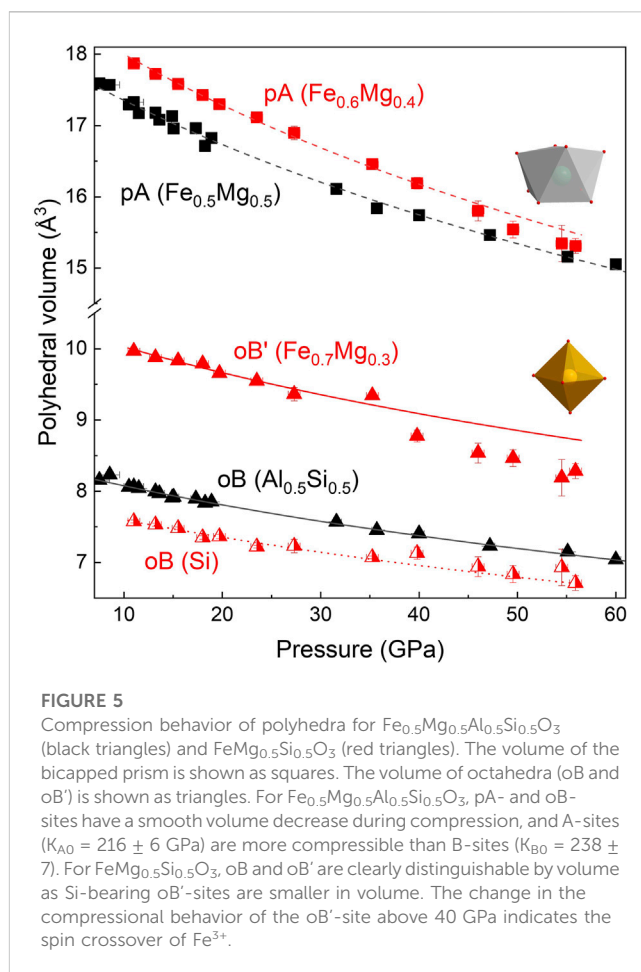
Synchrotron Mössbauer spectra were collected up to 95 GPa for $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ (Figure 2) and up to 62 GPa for

$\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ (Figure 3). The $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ sample contains noticeable amounts of ferrous iron. Due to the broadness of the doublet attributed to pA-site Fe^{2+} , the precise determination of Fe^{3+} content at ambient conditions is difficult.



However, from the Mössbauer spectra collected at high pressure, we can estimate the Fe^{2+} content to be approximately 16(4)%. This suggests the presence of less than 1% oxygen vacancies, which, however, cannot be detected directly from our microprobe data (uncertainty of the oxygen content is $\sim 3\%$). Mössbauer spectroscopy at ambient conditions and at high pressures shows that all iron in $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ is Fe^{3+} (Supplementary Figure S1). Upon compression, Fe^{3+} in $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ does not undergo spin transition (Figure 6).

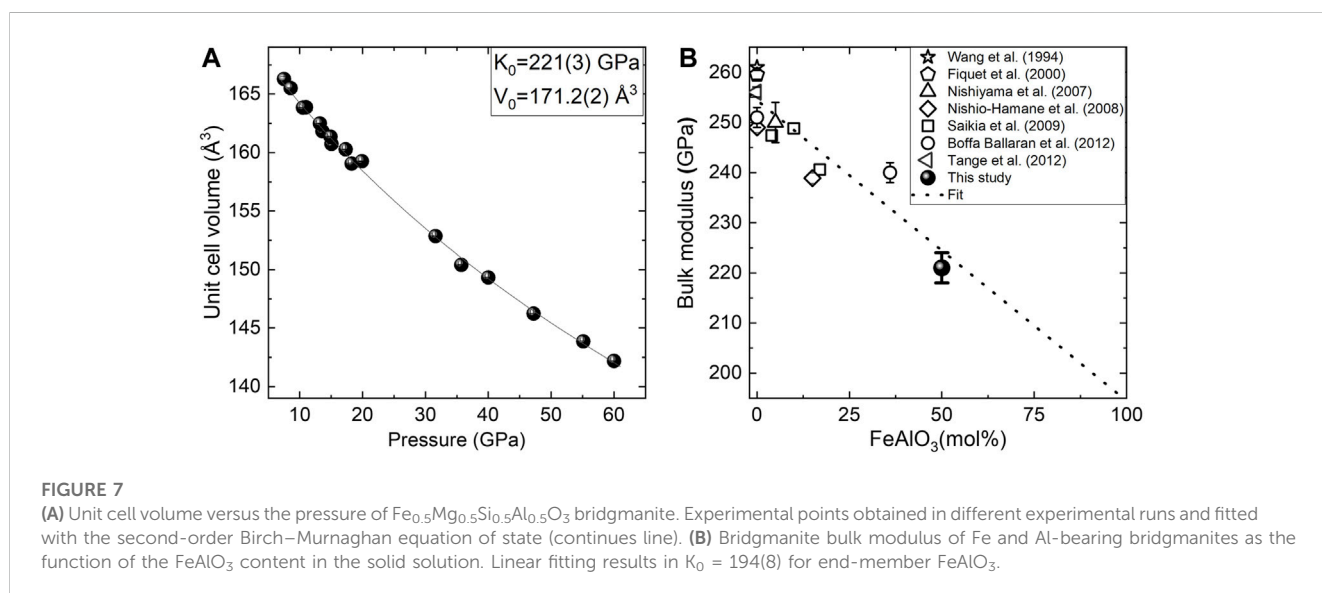
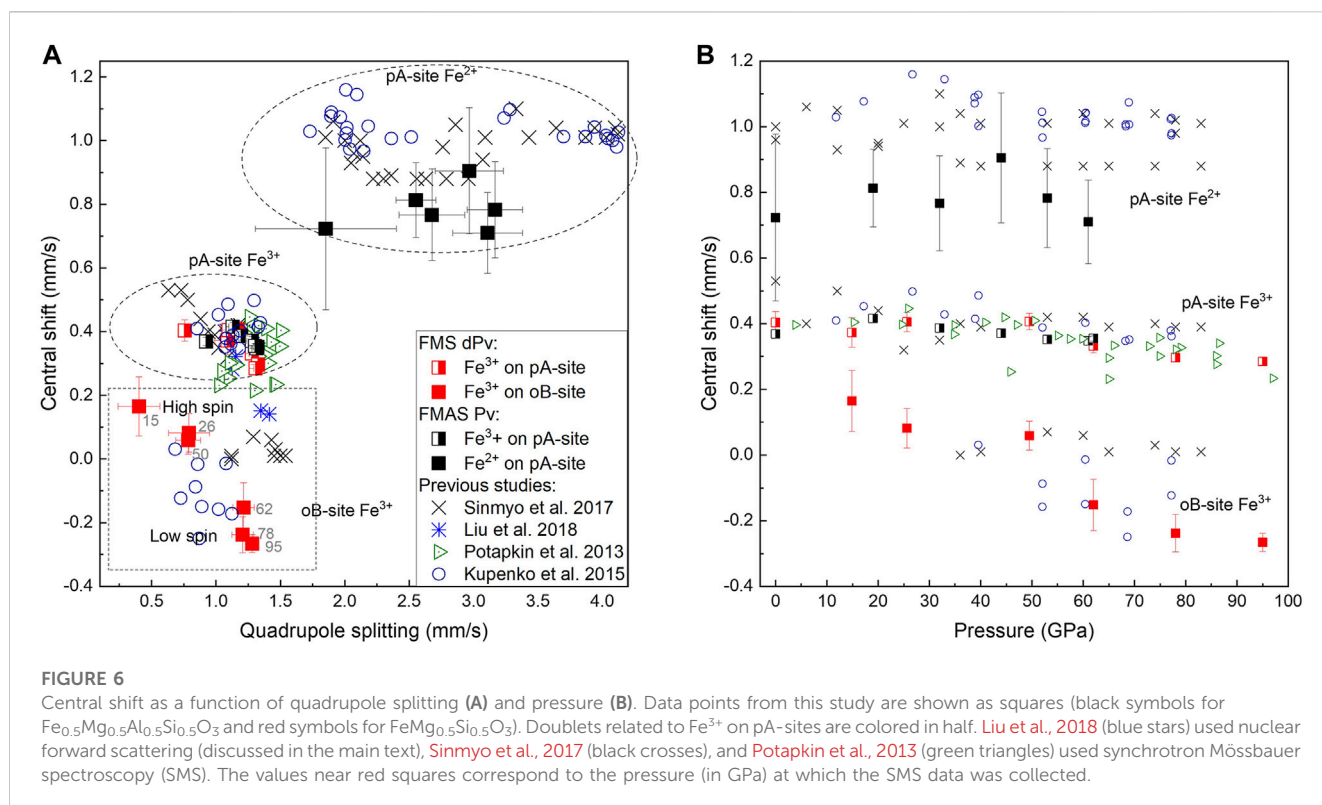
The near absence of Fe^{2+} in $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ together with the presence of high amounts of Fe^{3+} on oB-sites allowed us to unambiguously distinguish two doublets on Mössbauer spectra, which correspond to Fe^{3+} located on pA- and oB'-crystallographic sites. Even at pressures lower than the expected range for a spin crossover, one could still clearly observe the doublet for Fe^{3+} on oB'-sites (Figure 2). We show that the “New component” that appears in Sinmyo et al. (2017), and which was assigned to low spin Fe^{3+} , existed at



pressures before the spin crossover but was difficult to observe because of strong overlapping with intense A-site Fe^{2+} and Fe^{3+} doublets. Other evidence for this is the decrease in the Fe^{3+} doublet relative area, as reported by Sinmyo et al. (2017). Furthermore, hyperfine parameters that we associate with Fe^{3+} on oB-sites in $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ double perovskites are similar to those for a non-magnetic doublet, reported by Kuppenko et al., 2019, collected on $\zeta\text{-Fe}_2\text{O}_3$, which has a strongly distorted perovskite-like structure (Bykova et al., 2013; Bykova et al., 2016). Our results are generally consistent with a result on the sample with a composition similar to $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ from Liu et al. (2018), where they reported the central shift (CS) between the high-spin and low-spin states in the order of 0.2 mm/s based on the nuclear forward scattering experiment, which, however, does not allow to unambiguously determine the actual CS values. In agreement with the previous studies, we observe the QS increase with pressure for both studied compositions (Figure 6) (Sinmyo et al., 2017; Xiao et al., 2017; Supplementary Table S4). For example, doublets that correspond to Fe^{2+} on the pA-site in $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ perovskites have QS values of 2.50(5) mm/s at 19 GPa and 3.73(5) mm/s at 62 GPa (Figure 6).

3.4 Compressibility and effect of Al^{3+} and Fe^{3+} on the volume of the octahedra

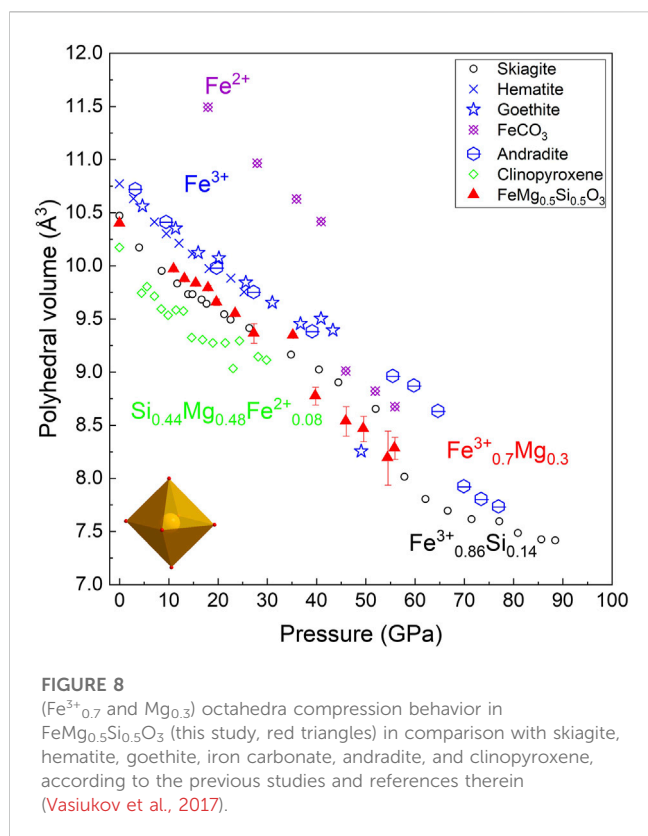
The volume per formula unit, as a function of pressure at ambient temperature, is shown for the two samples in Figure 4.



Discontinuities corresponding to the phase transition from the LiNbO_3 -type structure to a perovskite structure in the case of $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ and to phase transition from a new corundum-related structure to the double-perovskite structure in the case of $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ are observed at 8 ± 2 and 12 ± 2 GPa, respectively. The fitting of pressure–volume data for perovskite-structured phases with the second-order Birch–Murnaghan equation of state (BM2–EoS) resulted in a bulk modulus of 212(3) GPa for $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ and a bulk modulus of

199(6) GPa for $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$. The bulk moduli of the two perovskites are smaller than that of the MgSiO_3 perovskite (251 GPa, Ballaran et al., 2012).

In the case of $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$, we observed an increase in compressibility at pressures above 40 GPa, which is further supported by stress versus the Eulerian strain plot from the inset of Figure 4. The same observation was made by Liu et al. (2018) on a sample with a similar composition. The high-pressure single-crystal X-ray diffraction data allowed us to track individual polyhedral volumes of the studied



materials with increasing pressure. Although there is no change in the compressional behavior of pA- and oB-sites in the Al-rich sample, our experimental results clearly demonstrate the softening of Fe^{3+} -bearing B' octahedra in the $\text{FeMg}_{0.5}\text{Si}_{0.54}\text{O}_3$ sample (Figure 5). There is a very weak tendency in the volume increase of SiO_6 (B-site) octahedra with pressure across the spin crossover in oB'. A similar effect was observed in siderite (i.e., decrease in the volume of Fe^{2+}O_6 octahedra due to the spin transition and increase in C–O distances in CO_3 groups (Lavina et al., 2010)).

Our data indicate that in the case of $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ perovskites, a pA-site is more compressible than oB, as it was suggested previously (Sinmyo et al., 2017). When our result is considered together with previous data on the bulk compressibility of Fe-rich perovskites (Saikia et al., 2009; Ballaran et al., 2012; Dorfman et al., 2012; Glazyrin et al., 2014; Ismailova et al., 2016; Liu et al., 2018), one could see that it follows the general trend for the bulk modulus decrease with an increase in the Fe^{3+} content (Supplementary Figure S5). Moreover, the calculation of the bulk modulus for the Al-rich sample together with previously reported bulk moduli for Al, the Fe-bearing bridgmanite, allows us to constrain the bulk modulus of the pure FeAlO_3 end-member $K_0 = 194 \pm 8$ GPa (Figure 7). Potentially, one could estimate the composition from polyhedra volumes; however, this result might be affected by the uncertainty of Fe spin and oxidation states.

Although it is generally assumed that the volume collapse of Fe^{3+} -rich bridgmanite at pressures above 40 GPa is associated with the spin transition, our new data on the Fe^{3+} -bearing octahedra volume remain controversial. Figure 8 shows the oB' volumes of the $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ sample (Mg and Fe^{3+} -bearing octahedra) at all pressures are lower than that reported for pure Fe^{3+} octahedra in andradite, goethite, and hematite, despite the presence of Mg on the

oB'-site suggesting a volume increase compared to pure Fe^{3+} octahedra. The low spin state of Fe^{3+} at ambient pressures was previously reported for various metal–organic compounds (Nihei et al., 2007), so one of the possible explanations for the observed effect is the low spin state of octahedral Fe^{3+} at low pressure. Indeed, the estimated ionic radii of $\text{Fe}^{3+}_{0.7}\text{Mg}_{0.3}$, assuming a low spin state of ferric iron, is lower than that for the high spin state of Fe^{3+} .

4 Conclusion

We investigated the high-pressure crystal chemistry of well-characterized crystalline materials with two compositions, $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ and $\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$, synthesized by the Kawai-type multi-anvil apparatus. We performed a series of compressibility experiments in diamond anvil cells up to 95 GPa. During compression, we collected single-crystal X-ray diffraction patterns and Mössbauer spectra, which allowed us to follow up on the changes in the crystal chemistry and Fe spin state during compression.

$\text{Fe}_{0.5}\text{Mg}_{0.5}\text{Al}_{0.5}\text{Si}_{0.5}\text{O}_3$ with the LiNbO_3 -type structure at ambient conditions transforms into the perovskite-type phase at 8 GPa, and no spin transition was observed. On the other hand, $\text{FeMg}_{0.5}\text{Si}_{0.5}\text{O}_3$ has a novel structure at ambient conditions: the low-pressure phase is the corundum-related type (space group $R\bar{3}$), and at above approximately 12 GPa, it transforms into a new silicate double perovskite. An outstanding feature of the silicate double perovskite structure is having two individual octahedral sites: one occupied by Si only, and the other by iron and magnesium. Single-crystal X-ray diffraction Mössbauer spectroscopy data showed the compressibility changes of individual polyhedra and the variation in Mössbauer hyperfine parameters.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary Material; structural data is deposited in the CCDC database repository (<https://www.ccdc.cam.ac.uk/structures/>), accession numbers 2294965, 2294966, 2294967, 2294968. Further inquiries can be directed to the corresponding authors.

Author contributions

IK: formal analysis, investigation, writing—original draft, and writing—review and editing. BW, EK, TI, ZL, CM, ArC, TK, MH, and AIC: writing—review and editing. LD: conceptualization, supervision, and writing—original draft.

Funding

The authors declare that financial support was received for the research, authorship, and/or publication of this article. This work is supported by the Advanced Grant of the European Research Council (ERC) under the Horizon 2020 research and innovation program of the European Union (No. 787527) to TK.

Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Publisher's note

All claims expressed in this article are solely those of the authors and do not necessarily represent those of their affiliated

organizations, or those of the publisher, the editors, and the reviewers. Any product that may be evaluated in this article, or claim that may be made by its manufacturer, is not guaranteed or endorsed by the publisher.

Supplementary material

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fchem.2023.1258389/full#supplementary-material>

References

- Agilent (2014). *CrysAlis^{pro} data collection and processing software for agilent X-ray diffractometers*. Yarnton, Oxfordshire, England: Agilent Technologies Ltd.
- Andraut, D., Bolfan-Casanova, N., and Guignot, N. (2001). Equation of state of lower mantle (Al,Fe)-MgSiO₃ perovskite. *Earth Planet. Sci. Lett.* 193, 501–508. doi:10.1016/S0012-821X(01)00506-4
- Ballaran, T. B., Kurnosov, A., Glazyrin, K., Frost, D. J., Merlini, M., Hanfland, M., et al. (2012). Effect of chemistry on the compressibility of silicate perovskite in the lower mantle. *Earth Planet. Sci. Lett.* 333–334, 181–190. doi:10.1016/j.epsl.2012.03.029
- Bykova, E., Bykov, M., Prakapenka, V., Konôpková, Z., Liermann, H. P., Dubrovinskaia, N., et al. (2013). Novel high pressure monoclinic Fe₂O₃ polymorph revealed by single-crystal synchrotron X-ray diffraction studies. *High. Press. Res.* 33, 534–545. doi:10.1080/08957959.2013.833613
- Bykova, E., Dubrovinsky, L., Dubrovinskaia, N., Bykov, M., McCammon, C., Ovsyannikov, S. V., et al. (2016). Structural complexity of simple Fe₂O₃ at high pressures and temperatures. *Nat. Commun.* 7, 10661. doi:10.1038/ncomms10661
- Dorfman, S. M., Shieh, S. R., Meng, Y., Prakapenka, V. B., and Duffy, T. S. (2012). Synthesis and equation of state of perovskites in the (Mg, Fe)3Al₂Si₃O₁₂ system to 177 GPa. *Earth Planet. Sci. Lett.* 357–358, 194–202. doi:10.1016/j.epsl.2012.09.024
- Fei, Y., Ricolleau, A., Frank, M., Mibe, K., Shen, G., and Prakapenka, V. (2007). Toward an internally consistent pressure scale. *Proc. Natl. Acad. Sci.* 104, 9182–9186. doi:10.1073/pnas.0609013104
- Fiquet, G., Dewaele, A., Andraut, D., Kunz, M., and Le Bihan, T. (2000). Thermoelastic properties and crystal structure of MgSiO₃ perovskite at lower mantle pressure and temperature conditions. *Geophys. Res. Lett.* 27, 21–24. doi:10.1029/1999GL008397
- Frost, D. J., Liebske, C., Langenhorst, F., McCammon, C., Tronnes, R. G., and Rubie, D. C. (2004). Experimental evidence for the existence of iron-rich metal in the Earth's lower mantle. *Nature* 428, 409–412. doi:10.1038/nature02413
- Glazyrin, K., Boffa Ballaran, T., Frost, D. J., McCammon, C., Kantor, A., Merlini, M., et al. (2014). Magnesium silicate perovskite and effect of iron oxidation state on its bulk sound velocity at the conditions of the lower mantle. *Earth Planet. Sci. Lett.* 393, 182–186. doi:10.1016/j.epsl.2014.01.056
- Ishii, T., Liu, Z., and Katsura, T. (2019). A breakthrough in pressure generation by a kawai-type multi-anvil apparatus with tungsten carbide anvils. *Engineering* 5, 434–440. doi:10.1016/j.eng.2019.01.013
- Ishii, T., Shi, L., Huang, R., Tsujino, N., Druzhbin, D., Myhill, R., et al. (2016). Generation of pressures over 40 GPa using Kawai-type multi-anvil press with tungsten carbide anvils. *Rev. Sci. Instrum.* 87, 024501–024508. doi:10.1063/1.4941716
- Ishii, T., Sinmyo, R., Komabayashi, T., Ballaran, T. B., Kawazoe, T., Miyajima, N., et al. (2017). Synthesis and crystal structure of LiNbO₃-type Mg₅Al₂Si₃O₁₂: A possible indicator of shock conditions of meteorites. *Am. Mineralogist* 102, 1947–1952. doi:10.2138/am-2017-6027
- Ismailova, L., Bykova, E., Bykov, M., Cerantola, V., McCammon, C., Boffa Ballaran, T., et al. (2016). Stability of Fe, Al-bearing bridgmanite in the lower mantle and synthesis of pure Fe-bridgmanite. *Sci. Adv.* 2, e1600427. doi:10.1126/sciadv.1600427
- Kantor, I., Prakapenka, V., Kantor, A., Dera, P., Kurnosov, A., Sinogeikin, S., et al. (2013). BX90: A new diamond anvil cell design for X-ray diffraction and optical measurements. *Rev. Sci. Instrum.* 83, 125102. doi:10.1063/1.4768541
- Kubo, A., Yagi, T., Ono, S., and Akaogi, M. (2000). Compressibility of Mg_{0.9}Al_{0.2}Si_{0.9}O₃ perovskite. *Proc. Jpn. Acad.* 78, 103–107. doi:10.2183/pjab.76.103
- Kupenko, I., Aprilis, G., Vasiukov, D. M., McCammon, C., Chariton, S., Cerantola, V., et al. (2019). Magnetism in cold subducting slabs at mantle transition zone depths. *Nature* 570, 102–106. doi:10.1038/s41586-019-1254-8
- Kurnosov, A., Kantor, I., Boffa-Ballaran, T., Lindhardt, S., Dubrovinsky, L., Kuznetsov, A., et al. (2008). A novel gas-loading system for mechanically closing of various types of diamond anvil cells. *Rev. Sci. Instrum.* 79, 045110–045115. doi:10.1063/1.2902506
- Lavina, B., Dera, P., Downs, R. T., Yang, W., Sinogeikin, S., Meng, Y., et al. (2010). Structure of siderite FeCO₃ to 56 GPa and hysteresis of its spin-pairing transition. *Phys. Rev. B* 82, 064110. doi:10.1103/PhysRevB.82.064110
- Liu, J., Dorfman, S. M., Zhu, F., Li, J., Wang, Y., Zhang, D., et al. (2018). Valence and spin states of iron are invisible in Earth's lower mantle. *Nat. Commun.* 9, 1284–1289. doi:10.1038/s41467-018-03671-5
- Liu, Z., Dubrovinsky, L., McCammon, C., Ovsyannikov, S. V., Koemets, I., Chen, L., et al. (2019). A new (Mg_{0.5}Fe_{0.53+})(Si_{0.5}Al_{0.53+})O₃ LiNbO₃-type phase synthesized at lower mantle conditions. *Am. Mineralogist* 104, 1213–1216. doi:10.2138/am-2019-7070
- Mao, H. K., Hemley, R. J., Fei, Y., Shu, J. F., Chen, L. C., Jephcoat, A. P., et al. (1991). Effect of pressure, temperature, and composition on lattice parameters and density of (Fe,Mg)SiO₃ perovskites to 30 GPa. *J. Geophys. Res.* 96, 8069–8079. doi:10.1029/91JB00176
- Mao, H. K., Xu, J., and Bell, P. M. (1986). Calibration of the ruby pressure gauge to 800 kbar under quasi-hydrostatic conditions. *J. Geophys. Res.* 91, 4673–4676. doi:10.1029/jb091ib05p04673
- Mao, Z., Lin, J., Yang, J., Inoue, T., and Prakapenka, V. B. (2015). Effects of the Fe³⁺ spin transition on the equation of state of bridgmanite. *Geophys. Res. Lett.* 42 (11), 4335–4342. doi:10.1002/2015GL064400
- Nihei, M., Shiga, T., Maeda, Y., and Oshio, H. (2007). Spin crossover iron(III) complexes. *Coord. Chem. Rev.* 251, 2606–2621. doi:10.1016/j.ccr.2007.08.007
- Nishiyama, N., Yagi, T., Ono, S., Gotou, H., Harada, T., and Kikegawa, T. (2007). Effect of incorporation of iron and aluminum on the thermoelastic properties of magnesium silicate perovskite. *Phys. Chem. Min.* 34, 131–143. doi:10.1007/s00269-006-0134-6
- Palatinus, L., and Chapuis, G. (2007). SUPERFLIP - a computer program for the solution of crystal structures by charge flipping in arbitrary dimensions. *J. Appl. Crystallogr.* 40, 786–790. doi:10.1107/S0021889807029238
- Petricek, V., Dušek, M., and Palatinus, L. (2014). Crystallographic computing system JANA2006: general features. *Z. für Kristallogr.* 229, 345–352. doi:10.1515/zkri-2014-1737
- Potapkin, V., Chumakov, A. I., Smirnov, G. V., Celse, J. P., Rüffer, R., McCammon, C., et al. (2012). The ⁵⁷Fe synchrotron Mössbauer Source at the ESRF. *J. Synchrotron Radiat.* 19, 559–569. doi:10.1107/S0909049512015579
- Potapkin, V., McCammon, C., Glazyrin, K., Kantor, A., Kupenko, I., Prescher, C., et al. (2013). Effect of iron oxidation state on the electrical conductivity of the Earth's lower mantle. *Nat. Commun.* 4, 1427–1429. doi:10.1038/ncomms2436
- Rüffer, R., and Chumakov, A. I. (1996). Nuclear resonance beamline at ESRF. *Hyperfine Interact.* 97, 589–604. doi:10.1007/BF02150199
- Saikia, A., Ballaran, T. B., and Frost, D. J. (2009). The effect of Fe and Al substitution on the compressibility of MgSiO₃-perovskite determined through single-crystal X-ray diffraction. *Phys. Earth Planet. Interiors* 173, 153–161. doi:10.1016/j.pepi.2008.11.006
- Shannon, R. D. (1976). Revised effective ionic radii and systematic studies of interatomic distances in halides and chalcogenides. *Acta Crystallogr. Sect. A* 32, 751–767. doi:10.1107/S0567739476001551
- Sheldrick, G. M. (2015). SHELXT - integrated space-group and crystal-structure determination. *Acta Crystallogr. A* 71, 3–8. doi:10.1107/S2053273314026370
- Shukla, G., and Wentzcovitch, R. M. (2016). Spin crossover in (Mg,Fe³⁺)(Si,Fe³⁺)O₃bridgmanite: effects of disorder, iron concentration, and temperature. *Phys. Earth Planet. Interiors* 260, 53–61. doi:10.1016/j.pepi.2016.09.003

- Sinmyo, R., McCammon, C., and Dubrovinsky, L. (2017). The spin state of Fe 3+ in lower mantle bridgmanite. *Am. Mineralogist* 102, 1263–1269. doi:10.2138/am-2017-5917
- Sinogeikin, S. V., Zhang, J., and Bass, J. D. (2004). Elasticity of single crystal and polycrystalline MgSiO₃ perovskite by Brillouin spectroscopy. *Geophys Res. Lett.* 31, 019559. doi:10.1029/2004gl019559
- Tsuchiya, T., Tsuchiya, J., Umemoto, K., and Wentzcovitch, R. M. (2004). Phase transition in MgSiO₃ perovskite in the earth's lower mantle. *Earth Planet Sci. Lett.* 224, 241–248. doi:10.1016/j.epsl.2004.05.017
- Vanpeteghem, C. B., Zhao, J., Angel, R. J., Ross, N. L., and Bolfan-Casanova, N. (2006). Crystal structure and equation of state of MgSiO₃ perovskite. *Geophys Res. Lett.* 33, L03306–L03313. doi:10.1029/2005GL024955
- Vasiukov, D. M., Ismailova, L., Kuppenko, I., Cerantola, V., Sinmyo, R., Glazyrin, K., et al. (2017). Sound velocities of skiaegite–iron–majorite solid solution to 56 GPa probed by nuclear inelastic scattering. *Phys. Chem. Min.* 2017, 397–404. doi:10.1007/s00269-017-0928-8
- Xiao, Y., Yang, J., Chow, P., Mao, Z., Tomioka, N., Lin, J.-F., et al. (2017). Equation of state and hyperfine parameters of high-spin bridgmanite in the Earth's lower mantle by synchrotron X-ray diffraction and Mössbauer spectroscopy. *Am. Mineralogist* 102, 357–368. doi:10.2138/am-2017-5770
- Zhu, F., Liu, J., Lai, X., Xiao, Y., Prakapenka, V., Bi, W., et al. (2020). Synthesis, elasticity, and spin state of an intermediate MgSiO₃-FeAlO₃ bridgmanite: implications for iron in earth's lower mantle. *J. Geophys. Res. Solid Earth* 125, 019964. doi:10.1029/2020JB019964