

# 1 **MagellanPlus Workshop: Mission-specific platform approaches to** 2 **assessing natural hazards that impact society**

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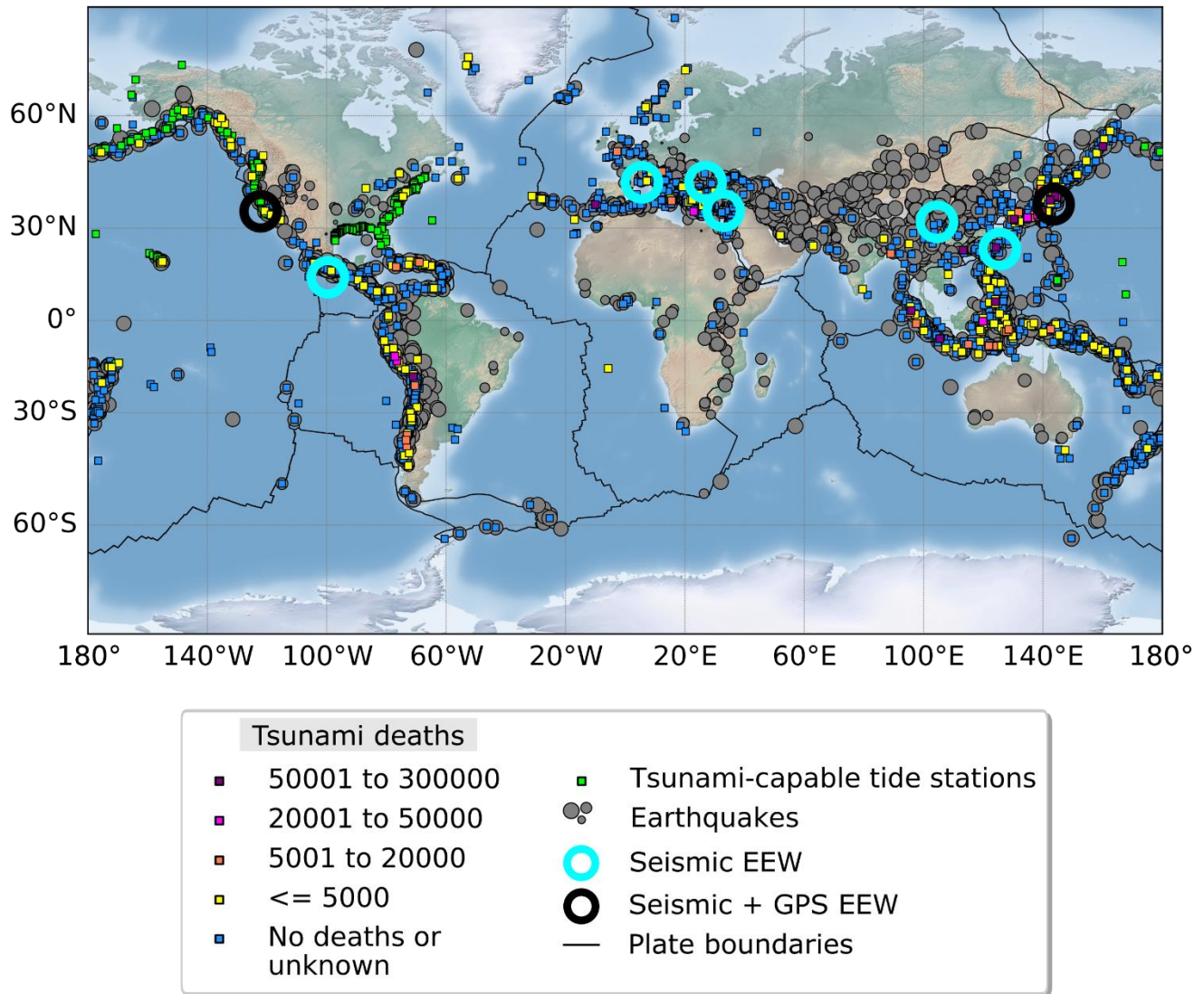
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13 **Abstract.** Oceanic natural hazards pose threats to coastal communities worldwide. These include earthquakes, tsunamis,  
14 submarine landslide, volcanic eruptions, and tropical cyclones. Scientific ocean drilling can contribute to our understanding  
15 and assessment of these hazards through rapid response measurements of hazardous events, learning from past hazard records,  
16 and seafloor monitoring and observation. With the impending retirement of the *D/V JOIDES Resolution* and operational  
17 limitations of the *D/V Chikyu*, it is important to consider other options for achieving scientific ocean drilling goals. We  
18 convened a workshop in Lisbon, Portugal in July 2022 to identify locations where natural hazards, or preferably several  
19 different hazards, can be addressed with mission-specific platform (MSP) drilling, with consideration of further location-based  
20 workshops. Participants split into 3 working groups to develop hypotheses surrounding climate and tropical cyclones, slope  
21 failure, and processes at active margins that can be tested with MSP drilling and can be addressed using the unique capabilities  
22 of these platforms. We produced 13 questions or hypotheses with recommendations on specific areas or locations for drilling.  
23 Our hope is that the results of this workshop will lay the groundwork for future preproposals.

## 24 **1 Introduction**

25 Natural hazards associated with the ocean, including earthquakes, tsunamis, submarine landslides, volcanic eruptions, and  
26 tropical cyclones, can have a direct impact on coastal populations, and even affect populations located far away from the coast  
27 (Koppers and Coggon, 2020). These hazards may interact, such as when tsunamis result in major direct damage and loss of  
28 life and also trigger submarine landslides, which themselves can produce tsunamis and damage subsea infrastructure like  
29 communications cables, oil and gas pipelines, and offshore wind turbines. In addition to these events, sea level rise and  
30 warming sea temperatures are resulting in more damaging tropical cyclones (Walsh et al., 2016), severe and nuisance coastal  
31 flooding (Morris and Renken, 2020; Vega et al., 2021), and larger-scale disruptions to ocean and atmospheric circulation

32 (Caesar et al., 2018). Tectonically and climatically driven hazards operate and interact over timescales that are societally  
 33 relevant, from seasonal to decadal (Telesca, 2007); and their records are preserved in the geologic record. Natural hazards can  
 34 be investigated in locations threatened by only one hazard, or locations where several different hazards are present. While  
 35 earthquakes and associated tsunamis more commonly originate on active margins, tsunamis triggered by



36  
 37 **Figure 1: Historic earthquakes 2150 B.C.E.-present (NGDC/WDS, 2023a), along with tsunami events 2000 B.C.E.-**  
 38 **present (NGCD/WDS, 2023b), tsunami-capable tide stations (NOS/CO-OPS, 2023), and earthquake early warning**  
 39 **(EEW) systems (Minson et al., 2015). Earthquakes shown meet at least one of the following criteria: caused at least US\$**  
 40 **1M in damage, resulted in 10 or more deaths, had a magnitude of 7.5 or more, had a Modified Mercalli intensity of X**  
 41 **or greater, or generated a tsunami,**

42

43 submarine landslides can occur on oceanic islands and passive margins as well (Fig. 1). Submarine landslides can themselves  
44 be triggered by earthquakes even on passive margins (Schulten et al., 2019), or through cyclic loading by large waves during  
45 tropical cyclones, which primarily affect the tropics but also the European Atlantic coast (Sainsbury et al., 2020). The  
46 prevalence of various hazards throughout geologic history could be investigated as well as linkages with longer-term trends in  
47 climate, ocean circulation, and tectonics. The paucity of warning systems for earthquakes and tsunamis (Fig. 1) demonstrates  
48 a particular opportunity in devising monitoring networks in other areas with high natural hazard prevalence.

49

50 The International Ocean Discovery Program (IODP) *2050 Science Framework* (Koppers and Coggon, 2020) lists investigating  
51 natural hazards impacting society as a strategic objective, with rapid response measurements of hazardous events, learning  
52 from past hazard records, and seafloor monitoring and observation identified as areas where scientific ocean drilling can  
53 contribute to understanding. Assessing earthquake and tsunami hazards is specifically identified as a flagship initiative in the  
54 *Framework*, while climate-related hazards fall under the flagship initiative to ground-truth future climate change. Mission-  
55 specific platforms (MSPs) can provide a significant advantage over large drillships in investigating natural hazards as they can  
56 potentially operate in shallower waters, more restricted environments, or in sea ice; MSPs can be specially tailored for  
57 deployment or monitoring of instrumentation; and they have the potential for more rapid deployment in response to new events,  
58 or repeat deployment over months or years to visit monitoring stations. They do have drawbacks as well, including limitations  
59 on water and coring depths as well as a restricted ability to perform shipboard analyses. MSPs are also usually smaller than  
60 the IODP vessels but are always be able to carry out the minimum of IODP measurements. MSPs therefore have promise for  
61 advancing scientific ocean drilling, but the circumstances in which they can be used requires definition.

62

63 MSPs are chosen by IODP to fulfill specific scientific objectives in certain conditions where the use of IODP vessels (*D/V*  
64 *JOIDES Resolution* or *D/V Chikyu*) is not appropriate. Whereas the *JOIDES Resolution* and *Chikyu* are dedicated drilling  
65 vessels, MSPs will be adapted to suit IODP needs as far as possible. Approaches for sample recovery of past MSPs include  
66 the use of wireline coring from geotechnical vessels or mining rigs on liftboats; and giant piston corer or seabed drills deployed  
67 from research vessels. There also exists the possibility to install borehole observatories, though this has not been tested on  
68 MSPs to date. The ECORD Science Operator (ESO) typically provides the necessary equipment for core curation and  
69 laboratory measurements onboard, usually in containerized facilities that are placed on-deck for the length of the cruise.  
70 Overall, the science party for an MSP expedition is oftentimes smaller than on an IODP vessel, and needs to be more flexible,  
71 as scheduling is more variable due to contracting of MSP vessels. After an MSP expedition, the majority of sampling and  
72 analysis takes place at the Onshore Science Party (OSP) Facility in Bremen, Germany, attended by the full science party.

73

74 MSP expeditions also aim at collecting *in situ* downhole logs. There are multiple approaches to collect downhole logs, with  
75 the standard one for MSP (and the *JOIDES Resolution*) being wireline logging, where tools are sent downhole connected to a

76 cable, a winch, and acquisition equipment. Wireline logging allows real time data acquisition against depth. The datasets  
77 acquired with this method on MSP expeditions can include borehole diameter, total and spectral gamma ray, formation  
78 resistivity or conductivity, magnetic susceptibility, acoustic velocity, borehole fluid properties and borehole wall images.  
79 As with MSP expeditions' flexibility for drilling platform and strategy, the logging equipment can be adapted to various  
80 operational setting and scientific objectives. For example, the European Petrophysics Consortium (EPC), part of ESO, has a  
81 set of wireline logging tools referred as "slimline" and designed by ALT, Mount Sopris, Antares and Geovista. These tools are  
82 relatively short (<2.5 m) lightweight (<20 kg) and narrow (4-6 cm diameter), and they can be run downhole as stand-alone  
83 sensors or combined into toolstrings of up to four tools and less than 40 kg. ESO-EPC provided wireline logging services on  
84 five MSP expeditions (see the methods section in Camoin et al. 2007; Mountain et al. 2010; Webster et al. 2011; Morgan et  
85 al. 2017; McNeill et al. 2019). Wireline logging services have also been provided by other service companies (see the methods  
86 section in Backman et al. 2006; Andr en et al. 2015). ESO also deployed logging-while-tripping slimline tools from seabed  
87 drills during Expedition 357 (see the methods section in Fr uh-Green et al., 2017).

88

89 With the impending retirement of the *D/V JOIDES Resolution* and operational limitations of the *D/V Chikyu*, it is important to  
90 consider other options for ocean drilling research. To explore the potential to use MSPs to address natural hazards impacting  
91 society, we convened a workshop in Lisbon, Portugal in July 2022 that brought together 28 researchers from around the world,  
92 including 9 early-career researchers, both in person and virtually (Daigle et al., 2023).

## 93 **2. Workshop Objectives**

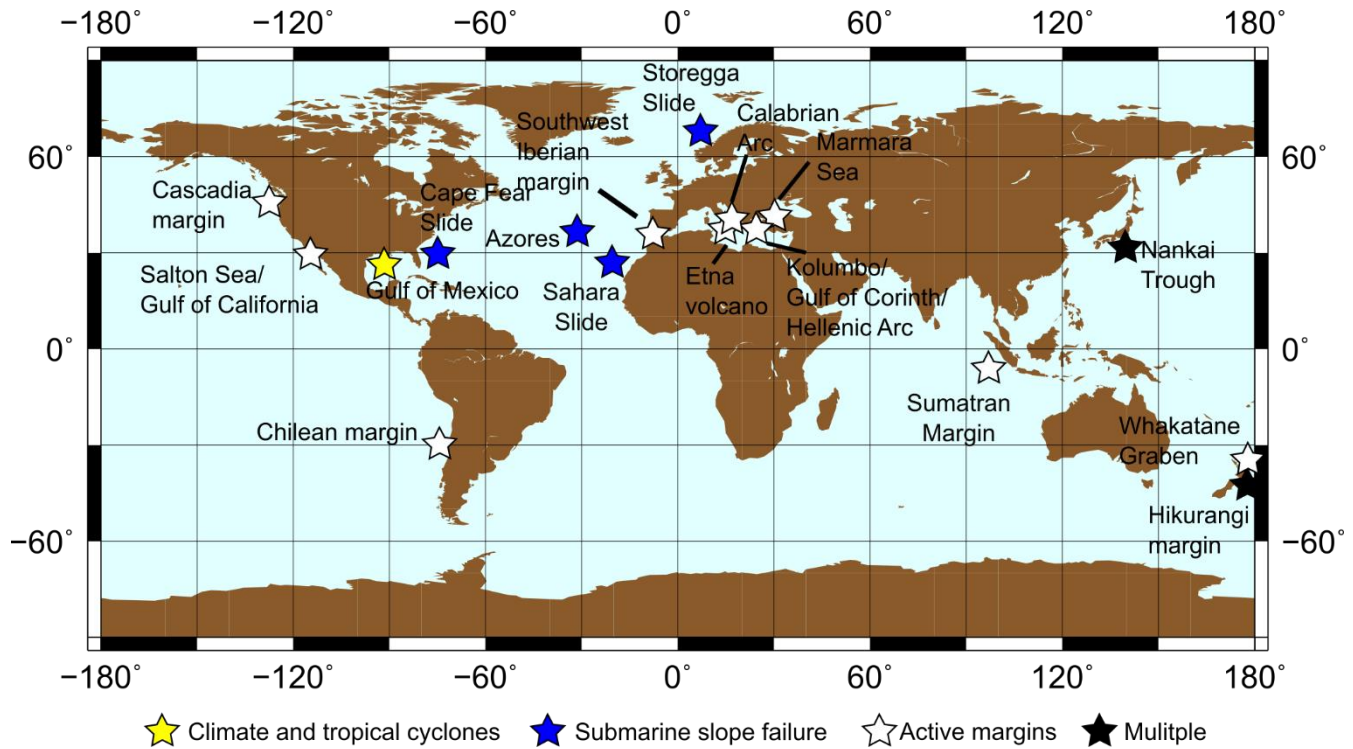
94 The objectives of the workshop were (1) to form working groups that would develop plans to address key scientific questions  
95 discussed at the meeting; (2) to identify locations where tsunamis, submarine landslides, volcanic eruptions, and tropical  
96 cyclones, or preferably several of these hazards, can be addressed with MSP drilling, with consideration of further location-  
97 based workshops; and (3) to develop a set of hypotheses that can be tested with MSP drilling which would lay the groundwork  
98 for future preproposals.

## 99 **3. Working Group Discussions**

100 During the workshop, we created three working groups that developed hypotheses and questions focused on three main topics:  
101 climate and tropical cyclones, slope failure, and processes at active margins. These hypotheses and questions can be used to  
102 develop drilling proposals for specific locations. Specific locations mentioned by the working groups are shown in Figure 2.

103 **3.1 Working Group 1: Climate and Tropical Cyclones**

104 The climate and tropical cyclones working group agreed that the northern Gulf of Mexico presents the best location to address  
105 questions about climate change and its influence on tropical cyclone frequency and intensity. The Gulf of Mexico is ideal



106  
107 **Figure 2. Locations recommended in the workshop color-coded by working group.**

108 because of its overall thick sediment column (up to 20 km; Galloway, 2008) and well-preserved sedimentary sections in salt-  
109 withdrawal minibasins, growth faults, river deltas, and basin-floor fans (Martin, 1978). The group phrased two hypotheses of  
110 global significance that may be addressed in the Gulf of Mexico.

111 **3.1.1 Hypothesis 1**

112 We hypothesize that past hurricane frequency and intensity in the Gulf of Mexico has responded to changes in the strength of  
113 ocean circulation patterns, most notably the strength of the Gulf Stream and the Atlantic Meridional Overturning Circulation  
114 (AMOC). Warm core eddies spun off the Loop Current can cause rapid intensification of hurricanes (e.g. Hurricane Katrina  
115 in 2005; Scharroo et al., 2005). Reconstruction of sea surface temperatures (SSTs) and storm frequency from sedimentary  
116 deposits in the Gulf of Mexico could be used to understand how much influence the variability in the Gulf Stream dynamics  
117 has on Gulf of Mexico storms.

118 *Drill site needs and potential outcomes.* MSPs could be used to collect very high-resolution storm and SST records in areas  
119 with continuous late Pleistocene to Holocene sedimentary sections. These time intervals contain many episodes of AMOC

120 slowdown (Heinrich Events, Younger Dryas) that would allow us to understand the impact of AMOC variability on hurricane  
121 development in this region.

122 *Potential drill sites.* Salt withdrawal minibasins along the Gulf of Mexico's margin contain suitable deposits, typically  
123 representing a sediment thickness of 100-200 m.

### 124 **3.1.2 Hypothesis 2**

125 Storm frequency has a strong influence on precipitation and streamflow in the western Gulf of Mexico and southwestern  
126 United States. Hurricane Harvey (2017) dropped more than 1500 mm of rain in parts of southeast Texas in less than a week  
127 and storm-related precipitation is expected to increase as climate warms over the 21st century (Bacmeister et al., 2018). Cores  
128 from the Gulf of Mexico have provided evidence for past occurrence of large flood events on the Mississippi River (e.g.,  
129 deglacial meltwater pulses; Clark et al., 1996; Brown and Kennett, 1998).

130 *Drill site needs and potential outcomes.* Drill sites should be located offshore of rivers that provide sedimentary flux from  
131 catchments that are affected by tropical cyclones. These locations can provide information on flooding events, particularly  
132 those during past climate intervals warmer than present day such as the mid-Holocene or previous interglacials. MSP drilling  
133 could include long cores, which could be combined with borehole or seafloor monitoring of development of mud flow events  
134 to establish modern depositional system analogs, and seismic surveys of the continental shelf to map potential sediment source  
135 areas.

136 *Potential drill sites.* Suitable locations can be found offshore the mouths of the Rio Grande, Colorado, Brazos, Sabine, and  
137 other rivers on the western Gulf of Mexico coast. Salt withdrawal minibasins offshore the mouth of the Rio Grande may make  
138 the best locations as they tend to preserve more continuous sedimentary records.

## 139 **3.2 Working Group 2: Submarine Slope Failure**

### 140 **3.2.1 Hypothesis 1: Preconditioning factors cause repeat landslides**

141 We hypothesize that through the Neogene, the presence of similar preconditioning factors has typically caused submarine  
142 landslides to occur repeatedly at the same location. Preconditioning factors can include sediment properties, pore fluid  
143 pressure, fluid chemistry, mechanical stratigraphy, and oversteepening/slope angle (Hampton et al., 1996). These can be  
144 investigated with hydraulic piston cores (equivalent to the *JOIDES Resolution* advanced piston coring (APC) tool) down to  
145 1000 m water depth in most cases. Required measurements include core physical properties and pore fluid chemistry, in situ  
146 pore pressure and heat flow measurements, and anything that can provide evidence of changes in stresses, e.g., borehole  
147 breakouts or drilling-induced tensile fractures obtained from multi-axis caliper or borehole image logs. In situ observatories  
148 can help establish trends in pore pressure, heat flow, and stress over time.

149 *Drill site needs and potential outcomes.* Drill sites should be located near the headscarp of a submarine landslide. Core analysis,  
150 downhole measurements, and in situ observatories will provide evidence of the presence or absence of preconditioning factors,  
151 and how these change over time.

152 *Potential drill sites.* There are many locations that could be targeted on both active and passive margins. The Cape Fear Slide,  
153 Storegga Slide, and Sahara Slide are potential targets on passive margins, while active margins such as the Nankai Trough, the  
154 Sumatran Margin, and Hikurangi Margin could be investigated.

### 155 **3.2.2 Hypothesis 2: Landslides are externally triggered**

156 Preconditioning factors are necessary but not sufficient for submarine landslides, and in most cases an external trigger is  
157 needed. Triggers include earthquakes, storm wave loading, abrupt shifts in sedimentation, tectonic oversteepening, changes in  
158 ocean temperature, gas hydrate dissociation, volcanism, and salt diapirism (e.g., Urlaub et al, 2013). This hypothesis can be  
159 tested by recovering cores, fluid samples, and pore pressures to demonstrate that slope failure could occur in an area that would  
160 otherwise be stable. Note that, if hypothesis 1 is true, then one initiation can cause the area to be prone to failure for many  
161 years afterwards.

162 *Drill site needs and potential outcomes.* Ideal sites to test this hypothesis should be locations with no evidence of slope failure,  
163 but near known submarine landslides. Measured properties could be used as inputs for models to predict conditions for slope  
164 failure.

165 *Potential drill sites.* This hypothesis could be tested at the same locations as the previous hypothesis, with the addition of  
166 volcanic islands such as the Azores.

### 167 **3.2.3 Hypothesis 3: Retrogressive failure leaves a signature**

168 Retrogressive failure leaves a detectable perturbation of pore pressure and heat flow. The pore pressure perturbation is related  
169 to failure propagation upslope, and the heat flow perturbation is a leftover signature of it. Anomalously high heat flow also  
170 could be evidence of salt diapirism. This technique can also be used to infer gas hydrate (in)stability. This hypothesis would  
171 best be tested with in situ measurements and longer-term monitoring with dense spatial sampling. One particular location that  
172 should be targeted is the distal toe of a submarine landslide, although these may be deeper than MSP capabilities (~1700 m)  
173 in most locations. For example, the toe of the Cape Fear Slide is at about 5.4 km water depth (Popenoe et al., 1993), while that  
174 of the Storegga Slide is at 3-3.5 km water depth (Bugge et al., 1988).

175 *Drill site needs and potential outcomes.* Potential sites should be located at submarine landslides with known retrogressive  
176 failure. Long-term measurements of pore pressure and heat flow can provide evidence of the current state of the subsurface  
177 around submarine landslides and its response to external perturbations.

178 *Potential drill sites.* The Cape Fear Slide is a well studied example of retrogressive failure that could be targeted to test this  
179 hypothesis.

#### 180 **3.2.4 Hypothesis 4: Active margins have more earthquakes than landslides**

181 Submarine landslides occur less frequently than earthquakes on active margins because of (a) seismic strengthening and (b)  
182 sediment accumulation rates. It is established that seismic shaking tends to increase the shear strength of sediments (Sawyer  
183 and DeVore, 2015), so slope failure on an active margin would tend to require a large amount of sediment to accumulate  
184 between earthquakes.

185 *Drill site needs and potential outcomes.* To test this hypothesis, high-resolution seismic data, core data to understand strength  
186 and sedimentation rate, and a good paleoseismological record are required. A drilling strategy to test this hypothesis would  
187 consist of a transect across an accretionary prism with perched basins that preserved mass transport deposits.

188 *Potential drill sites.* The Nankai Trough or Hikurangi Margin make excellent candidates to test this hypothesis due to the large  
189 volume of existing data.

190

### 191 **3.3 Working Group 3: Active Margins**

#### 192 **3.3.1 Hypothesis 1: Major earthquakes have precursors**

193 Transient events are observed at plate boundaries worldwide, but we currently do not understand how these relate to the timing  
194 of larger earthquakes. In a few cases, slow slip events at subduction zones have been observed in the weeks to months prior to  
195 great megathrust earthquakes (e.g., the Tohoku-Oki 2011 and Iquique 2014 earthquakes) (Ito et al., 2013; Ruiz et al., 2014).  
196 Resolving whether great earthquakes have precursors is a societally important question, critical for more effective, short-term  
197 earthquake forecasting. As most subduction zone earthquakes nucleate on offshore megathrusts, detailed offshore monitoring  
198 of potential transient deformation is required.

199 *Drill site needs and potential outcomes.* Drill sites should be located near active faults in areas with large earthquakes. Due to  
200 the high-noise ocean environment, the most viable tool to undertake this type of monitoring is borehole observatories,  
201 specifically detecting volumetric strain using changes in formation pore pressure as a proxy; as done in boreholes offshore  
202 subduction margins in Costa Rica, Japan, and New Zealand (Davis et al., 2015; Araki et al., 2017; Wallace et al., 2019).

203 *Potential drill sites.* Highly active transform faults that are the site of frequently recurring moderate to large magnitude  
204 earthquakes, such as the Gofar transform fault with quasi-periodic Mw 6-7 events every 5-6 years (McGuire, 2008), could be  
205 ideal targets to investigate potential precursors through installation of borehole observatories. Amphibious experiments  
206 complementing the already existing infrastructure could be developed in both the Chilean Margin (Barrientos et al., 2018) and  
207 the Marmara Sea (MARSite; Özel et al., 2017). The Salton Sea and the Gulf of California may also be potential sites, as a  
208 southward extension of the U.S. Geological Survey Parkfield, CA network, but are complicated by transtensional deformation  
209 (Brothers et al., 2009). In addition, the Hikurangi Margin, Hellenic Arc, and Southwest Iberian Margin could be other locations  
210 targeted to test this hypothesis.

211



### 213 3.3.2 Hypothesis 2: Rupture barriers are persistent

214 Several hypotheses exist to explain rupture terminations that appear to be persistent over the timescale of available earthquake  
215 records. These range from the presence of plate boundary transitions (Wallace et al., 2012), to changes in upper plate/lower  
216 plate properties and/or geometrical changes (Collot et al., 2004), stress heterogeneity (Huang, 2018), and the presence of  
217 positive relief structures such as seamounts or ridges (Sparkes et al., 2010). However, no conclusive evidence exists to confirm  
218 that these historical or recent rupture terminations are persistent over longer timescales. Potential boundary-breaking  
219 earthquake ruptures could lead to earthquakes that substantially exceed historical or instrumentally observed magnitudes  
220 (maybe even  $M > 10$ ) and that have recurrence intervals of several 1000s of years or more (Goldfinger et al., 2013). This is  
221 much longer than the current length of available paleoseismic records along any subduction zone around the world, meaning  
222 we cannot reliably exclude their occurrence.

223

224 To evaluate the persistence of rupture/segment barriers and unveil (or discard) the existence of these larger-than-observed  
225 earthquakes, the available paleoseismic records need to be extended back in time. Long-term paleoseismic records are thus  
226 needed along the entire length of subduction margins. This requires sedimentary records in both the onshore (coastal and lake)  
227 and offshore realm, and relies on the identification of secondary seismic effects such as shaking imprints (e.g. turbidites, soft-  
228 sediment deformation) and tsunami deposits. Identification of synchronous deposits along an entire subduction margin could  
229 hint towards these extreme events, but the current limitations of dating accuracy do not allow distinguishing single, large  
230 events from short-term successions of ruptures (e.g. stress triggering). Therefore, key records at locations that are currently  
231 believed to form rupture barriers are essential to verify the existence of imprint stacks, resulting from rupture cascades, while  
232 the presence of a single deposit could hint towards large, through-going ruptures.

233 *Drill site needs and potential outcomes.* Drill sites need to have historic earthquake records that show earthquake boundaries;  
234 a detailed understanding of sediment routing systems and several preferentially isolated depositional basins spanning the entire  
235 margin and containing high-resolution, continuous paleoseismic event deposit stratigraphies, for which sedimentary source  
236 areas can be well constrained to allow for reconstruction of past rupture areas; and a straight margin to avoid complications of  
237 changes in fault geometry (unless testing the role of plate margin geometry on arresting earthquake rupture is desired). The  
238 outcome of drilling would ideally be a long-term extension of the paleoseismic record.

239 *Potential drill sites.* The South American subduction zone is a potential target to investigate this hypothesis. It generated the  
240 largest known earthquake (magnitude 9.6, in 1960), and it is potentially able to generate the largest earthquakes of any  
241 subduction zone (Graham et al., 2021). In addition, many lakes along the margin could be sampled by continental drilling to  
242 constrain the paleoseismic record (Bernhardt et al., 2015). Other potential locations include the Marmara Sea, the Salton  
243 Sea/Gulf of California, Hikurangi Margin, Cascadia, the Hellenic Arc, and the Calabrian Arc.

### 244 **3.3.3 Hypothesis 3: Fault coupling characteristics are persistent in space and time**

245 Subduction plate interfaces commonly have spatially variable interseismic coupling, where strongly coupled segments are  
246 most likely to produce large earthquakes and weakly coupled segments tend to slip aseismically (Fagereng, 2011; Saito and  
247 Noda, 2022). It is unclear whether geodetically-constrained coupling patterns persist over multiple earthquake cycles, and also  
248 to what extent poorly coupled regions may slip during earthquakes that nucleate in adjacent coupled areas. Controls on coupling  
249 could include fluid pressure (Moreno et al., 2014); downgoing plate roughness (Ruff, 1989; Wang and Bilek, 2011; van  
250 Rijnsingen et al., 2018); or rock physical properties at depth, including grain size, mineralogy, and fluid composition (Chen et  
251 al., 2013; Scholz, 2019).

252 *Drill site needs and potential outcomes.* Direct evidence for temporal variation in locking may be sought from the paleoseismic  
253 record, by looking for evidence for past earthquakes, and how their spatial pattern compares with current geodetic locking.  
254 This would require a margin where the current geodetic locking pattern is well characterized, and where there are appropriate  
255 sites for paleoseismicity (e.g. Chilean margin). The expectation, if coupling characteristics are persistent in time and space,  
256 would be that in locked regions, there is evidence for large earthquakes (where the meaning of ‘large’ should be calculated  
257 from the size of the locked patch). These earthquakes should, however, not have propagated into poorly coupled segments.  
258 Another direct measure is variations in interseismic periods, if observatories can be installed and maintained over multiple  
259 earthquake cycles.

260 *Potential drill sites.* Sites with existing reflection seismic and other data that could be used to test this hypothesis include the  
261 Chilean Margin, Marmara Sea, the Hellenic Arc, Salton Sea/Gulf of California, Cascadia Nankai Trough, Hikurangi Margin,  
262 and Costa Rica Margin.

263

### 264 **3.3.4 Hypothesis 4: Subduction earthquakes are cyclic at multiple time scales**

265 The 2011 Mw 9.0 Tohoku-Oki Japan earthquake occurred in an area where scientists thought large magnitude earthquakes  
266 were less likely (Stein et al., 2012). In contrast, the 2010 Mw 8.8 Maule earthquake in Chile was somewhat expected, but  
267 extended beyond the boundaries of the seismic gap that was believed to be present (Métois et al., 2012). These recent, very  
268 high-magnitude earthquakes painfully highlight how little we still know about megathrust earthquake recurrence along  
269 subduction zones, despite recent advances and widespread paleoseismological studies. The concepts of seismic gaps and  
270 characteristic earthquakes are simplified, and need further refining.

271 Mapping the spatiotemporal behavior of megathrust earthquakes thus forms a crucial step towards validation of physics-based  
272 earthquake cycle models, but is currently not possible on sufficiently long timescales and/or spatial extents along any of the  
273 subduction zones. Instrumental and historical records are too short, and coastal records (tsunami deposits, uplifted terraces or  
274 corals, subsided paleosols) are affected by global eustatic sea-level change and do not extend far beyond the last maximum  
275 sea level high stand of the Holocene.

276 *Drill site needs and potential outcomes.* Scientific ocean (and continental) drilling and coring of high-resolution marine and/or  
277 lacustrine paleoseismic archives extending well into the Late Pleistocene and further back in time as the only reasonable  
278 approach to potentially deliver observational data on time-scales long enough to robustly test the earthquake (super-)cycle  
279 hypothesis. Therefore, drill sites need to have well preserved, continuous sedimentary sections.

280 *Potential drill sites.* Drilling to test this hypothesis could be conducted at the Chilean Margin, the Hikurangi Margin, and  
281 Cascadia.

282

### 283 **3.3.5 Hypothesis 5: Fault slip rates vary over multiple seismic cycles**

284 Fault slip rates are critical for seismic hazard assessment and can be calculated over a range of different time scales from < 1  
285 yr to millions of years. Variations between short-term and long-term slip rates have been recorded, bringing into question the  
286 usefulness of slip rates calculated over a particular time period for seismic hazard assessment, i.e., are geological rates relevant  
287 to apply to modern seismic hazard estimates? Do modern satellite derived slip rates (from GPS) give a good indication of the  
288 areas most at risk from earthquakes (e.g., Bell et al. 2011, Cowie et al. 2012; Fagereng and Biggs, 2019)? Our understanding  
289 of fault growth and fault slip is least well-constrained in the  $10^4$ - $10^6$  yr range (reviewed in Pan et al. 2022). There are a number  
290 of ways in which we can assess slip rates in the range of  $10^4$ - $10^6$  yrs using offshore MSP drilling/piston core data, including  
291 targeted giant piston coring or short drilled sections on either side of a fault (or one side of the fault only if horizons can be  
292 confidently correlated) to calculate offset and slip rate; using submarine paleoseismology to identify evidence of individual  
293 earthquake slip events in high resolution seismic data with ground truthing of age and timing; or an onshore-offshore approach  
294 for comparison of slip rates over different timescales on an active normal fault.

295 *Drill site needs and potential outcomes.* Potential locations to study slip rates through time would ideally have: i) high  
296 sedimentation rates, ii) high slip rates (to give the greatest chance of slip rate variations being resolved), iii) a constrained  
297 onshore record of uplift rates to compare with results of ocean/lake drilling, and iv) high-resolution seismic reflection data  
298 imaging basin stratigraphy in the hanging wall and ideally also in the footwall.

299 *Potential drill sites.* Drilling to test this hypothesis could be conducted at active rift zones, and potentially at suitable transform  
300 zones and subduction margins, e.g., Corinth Rift (Greece), Salton Sea/Gulf of California, Marmara Sea, Chilean margin,  
301 Hikurangi margin.

302

### 303 **3.3.6 Hypothesis 6: Faults grow rapidly to their full length**

304 Seismological and geodetic data reveal that earthquake rupture patterns are complex and variable, often with temporal and  
305 spatial clustering of events (e.g., Wells and Coppersmith, 1994). In contrast, ancient normal faults commonly observed in high-  
306 resolution 3D seismic reflection datasets reveal strikingly consistent patterns of displacement accumulation along strike,  
307 commonly described as the classic ‘bell shaped’ cumulative displacement profile with greatest overall displacement in the

308 center of the fault decreasing toward the tips (e.g. Cowie and Scholz 1992). Other mature faults within extended systems have  
309 consistent slip rates along strike, potentially indicating linkage of segments at depth. Exactly how faults evolve in terms of  
310 how multiple earthquake cycles and aseismic slip accrue to produce the long-term fault geometry is unknown. Most work into  
311 this problem has focused on normal faults, as they are associated with a convenient marker of fault growth in the form of  
312 sediment thickness increases in the hanging wall when a fault is active at the Earth's surface (i.e. it is a 'growth fault').  
313 Therefore, to address the question of "how faults grow" will likely require observations from active continental rifts.  
314 Understanding how faults develop has relevance to earthquake hazards because of how we interpret fault slip rates over  
315 different time periods (see section 3.3.5 above) and how earthquake slip builds up to fault slip over longer timescales. One  
316 category of normal fault relatively understudied is outer rise normal faults at subduction zones. These faults, caused by flexure  
317 of the downbending subducting plate, are a significant tsunami hazard and they are potential sites for investigating the normal  
318 fault evolution of this hypothesis.

319 *Drill site needs and potential outcomes.* It may be possible to investigate in detail how faults establish themselves and evolve  
320 both laterally and in terms of displacement accrual by identifying a study location where a very young fault exists at a shallow  
321 depth, which is well imaged by 3D or pseudo-3D high-resolution seismic reflection data. If age-constraints are available from  
322 drilling/piston coring, interpretation of high-resolution seismic reflection data will allow variations in sediment thickness  
323 between the hanging wall and footwall to be investigated at the millennial scale, and hence how slip has accrued along different  
324 parts of the fault. This timescale links the supramillennial scale of modern geological and seismological observation, and the  
325 million-year averaged observations from seismic reflection data in rifts.

326 *Potential drill sites.* Drilling to test this hypothesis could be conducted at the following locations: Corinth Rift (Greece)  
327 (McNeill et al., 2019; Nixon et al., 2016); Whakatane Graben (New Zealand) (Taylor et al., 2004); and outer rise normal fault  
328 systems at subduction margins.

329

### 330 **3.3.7 Hypothesis 7: Tectonic activity is linked to the timing of volcanic eruptions**

331 Active submarine volcanoes are linked to several of the deadliest geohazards such as tsunamis, earthquakes, landslides,  
332 pyroclastic material, and gas release. However, they pose difficult drilling conditions as they are covered by thin microbial  
333 crusts, and they have active hydrothermal systems and unstable slopes. MSP drilling presents an ideal way to sample these  
334 locations. Volcanic activity (frequency, volume, magma flux) is affected by tectonic activity at large length scales (e.g., plate  
335 tectonics) and smaller scales (e.g., development of preferential pathways for magma ascent due to local deformation). On the  
336 other hand, tectonic activity can also be affected by volcanism, for example in the case of magma ascent triggering localized  
337 seismic activity. This interconnection can be investigated through a high-resolution record of eruptive styles at individual  
338 centers that will give a characterization of their evolution in time, intensity, and spatial and temporal distribution. Large tectonic  
339 events recorded in sedimentary basins, like onlap surfaces, seismogenic turbidites, and homogenites, can correlate with  
340 activation/deactivation of different volcanic centers, changes in eruptive style, and particularly large explosive eruptions.

341 *Drill site needs and potential outcomes.* Drilling should be conducted at sites with a well preserved record of volcanism and  
342 tectonic events. Integration of paleoseismic and tephrochronological records need a good correlation between onshore and  
343 offshore records. Offshore geodesy would also be a priority to reconstruct the evolution of deformation around and within  
344 these volcanic centers. Possible correlations also exist with the paleoclimate record.  
345 *Potential drill sites.* This hypothesis could be tested by drilling at the Chilean Margin, Kolumbo submarine volcano northeast  
346 of Santorini (building on recent IODP drilling), Etna volcano, the Hikurangi Margin, the Hellenic Arc, and the Calabrian Arc.

#### 347 **4. Conclusions**

348 In this workshop, we defined questions and hypotheses about natural hazards that could be tested specifically with MSPs. The  
349 unique characteristics of MSPs compared to IODP drill ships require some shift in our thinking about how to interrogate the  
350 subsurface, but also present opportunities in terms of long-term monitoring, drilling in shallow water, and amphibious  
351 proposals. The workshop discussions recognized the growing needs of highly populated regions along tectonically active or  
352 hazardous settings for better assessment of earthquake, tsunami, volcanic, and climate hazards. A number of key  
353 questions/hypotheses, in combination with potential target locations and research strategies were formulated that urgently need  
354 to be tested in order to mitigate future risks.

355

356 MSP missions have the potential to contribute to the development of amphibious observatories, taking advantage of well-  
357 monitored onshore regions and leveraging partnerships with the continental drilling community. We considered that this  
358 strategy should be prioritized in the near future. The flexibility of an MSP and the relatively low costs in some cases make it  
359 worth the community effort in building a long-term infrastructure that allows us to monitor diverse geohazards from such  
360 observatories. The climate and tropical cyclones working group identified the Gulf of Mexico where key questions can be  
361 answered by MSP drilling. The submarine slope failure working group identified areas on both passive and active margins  
362 where hypotheses can be tested, including the Cape Fear Slide, the Storegga Slide, the Sahara Slide; the Nankai Trough, the  
363 Sumatran Margin, and the Hikurangi Margin; and volcanic islands such as the Azores. Finally, the active margins working  
364 group selected priority sites for the near future including the Chilean margin, the Hikurangi margin, the Marmara Sea, the  
365 Salton Sea/Gulf of California, Corinth rift, outer-rise normal fault systems. Other sites with high hazard exposure on active  
366 margins include regions around the Hellenic and Calabrian arcs, the Southwest Iberian, Cascadia, Sumatra, and Hikurangi  
367 margins. We hope that this work serves as the foundation for future drilling proposals in the regions we identified, or in other  
368 areas around the world.

369 **Data availability**

370 The datasets containing the historic earthquakes 2150 B.C.E-present (NGDC/WGS, 2023a), the tsunami events 2000 B.C.E.-  
371 present (NGDC/WGS, 2023b) and the tsunami-capable tide stations (NOS/CO-OPS, 2023) are available at their corresponding  
372 websites of the NOAA National Geophysical Data Center / World Data Service.

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390 **Competing interests**

391 The authors declare that they have no conflict of interest.

392 **Disclaimer**

393 Any use of trade, firm, or product names is for descriptive purposes only and does not imply endorsement by the U.S.  
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399 **References**

- 400 Andrén, T., Jørgensen, B. B., Cotterill, C., and the Expedition 347 Scientists: Expedition 347 summary, *Proc. IODP*, 347, 1-  
401 66, doi:10.2204/iodp.proc.347.101.2015, 2015.
- 402 Araki, E., Saffer, D. M., Kopf, A. J., Wallace, L. M., Kimura, T., Machida, Y., Ide, S., Davis, E., and the Expedition 365  
403 Scientists: Recurring and triggered slow-slip events near the trench at the Nankai Trough subduction megathrust, *Science*,  
404 356(6343), 1157-1160, doi:10.1126/science.aan3120, 2017.
- 405 Backman, J., Moran, K., McInroy, D. B., Mayer, L. A., and the Expedition 302 Scientists: Expedition 302 summary, *Proc.*  
406 *IODP*, 302, 1-22, doi:10.2204/iodp.proc.302.101.2006, 2006
- 407 Bacmeister, J. T., Reed, K. A., Hannay, C., Lawrence, P., Bates, S., Truesdale, J. E., Rosenbloom, N., and Levy, M.: Projected  
408 changes in tropical cyclone activity under future warming scenarios using a high-resolution climate model, *Climatic Change*,  
409 146, 547-560, doi:10.1007/s10584-016-1750-x, 2018.
- 410 Barrientos, S., and the National Seismological Center (CSN) Team: The seismic network of Chile, *Seismol. Res. Lett.*, 89(2A),  
411 467-474, doi:10.1785/0220160195, 2018.
- 412 Bell, R. E., McNeill, L. C., Henstock, T. J., and Bull, J. M.: Comparing extension on multiple time and depth scales in the  
413 Corinth Rift, Central Greece, *Geophys. J. Int.*, 186(2), 463-470, doi:10.1111/j.1365-246X.2011.05077.x, 2011.
- 414 Bernhardt, A., Melnick, D., Hebbeln, D., Lückge, A., and Strecker, M. R.: Turbidite paleoseismology along the active  
415 continental margin of Chile – Feasible or not?, *Quaternary Sci. Rev.*, 120, 71-92, doi:10.1016/j.quascirev.2015.04.001, 2015.
- 416 Brothers, D. S., Driscoll, N. W., Kent, G. M., Harding, A. J., Babcock, J. M., and Baskin, R. L.: Tectonic evolution of the  
417 Salton Sea inferred from seismic reflection data, *Nature Geosci.*, 2, 581-584, doi:10.1038/ngeo590, 2009.
- 418 Brown, P. A., and Kennett, J. P.: Megaflood erosion and meltwater plumbing changes during last North American deglaciation  
419 recorded in Gulf of Mexico sediments, *Geology*, 26(7), 599-602, doi:10.1130/0091-  
420 7613(1998)026<0599:MEAMPC>2.3.CO;2, 1998.
- 421 Bugge, T., Belderson, R. H., and Kenyon, N. H.: The Storegga Slide, *Philos. T. R. Soc. S-A*, 325(1586), 357-388,  
422 doi:10.1098/rsta.1988.0055, 1988.
- 423 Caesar, L., Rahmstorf, S., Robinson, A., Feulner, G., and Saba, V.: Observed fingerprint of a weakening Atlantic Ocean  
424 overturning circulation, *Nature*, 556, 191-196, doi:10.1038/s41586-018-0006-5, 2018.
- 425 Camoin, G. F., Iryu, Y., McInroy, D. B., and the Expedition 310 Scientists: Methods, *Proc. IODP*, 310, 1-43,  
426 doi:10.2204/iodp.proc.310.103.2007, 2007.

427 Chen, W. P., Yu, C. Q., Tseng, T. L., Yang, Z., Wang, C. yuen, Ning, J., and Leonard, T.: Moho, seismogenesis, and rheology  
428 of the lithosphere, *Tectonophys.*, 609, 491–503, doi:10.1016/j.tecto.2012.12.019, 2013.

429 Clark, P. U., Alley, R. B., Keigwin, L. D., Licciardi, J. M., Johnsen, S. J., and Wang, H.: Origin of the first global meltwater  
430 pulse following the Last Glacial Maximum, *Paleoceanogr. Paleoclimatology*, 11(5), 563-577, doi:10.1029/96PA01419, 1996.

431 Collot, J.-Y., Marcaillou, B., Sage, F., Michaud, F., Agudelo, W., Charvis, P., Graindorge, D., Gutscher, M.-A., and Spence,  
432 G.: Are rupture zone limits of great subduction earthquakes controlled by upper plate structures? Evidence from multichannel  
433 seismic reflection data acquired across the northern Ecuador–southwest Colombia margin, *JGR Solid Earth*, 109(B11),  
434 B11103, doi:10.1029/2004JB003060, 2004.

435 Cowie, P. A., and Scholz, C. H.: Displacement-length scaling relationship for faults: data synthesis and discussion, *J. Struct.*  
436 *Geol.*, 14(10), 1149-1156, doi:10.1016/0191-8141(92)90066-6, 1992.

437 Cowie, P. A., Roberts, G. P., Bull, J. M., and Visini, F.: Relationships between fault geometry, slip rate variability and  
438 earthquake recurrence in extensional settings, *Geophys. J. Int.*, 189(1), 143-160, doi:10.1111/j.1365-246X.2012.05378.x,  
439 2012.

440 Daigle., H., et al.: MagellanPlus Workshop: Mission-specific platform approaches to assessing natural hazards that impact  
441 society, MagellanPlus workshop report, ECORD/ICDP, <https://www.ecord.org/?ddownload=16319>, 2022.

442 Davis, E. E., Villinger, H., and Sun, T.: Slow and delayed deformation and uplift of the outermost subduction prism following  
443 ETS and seismogenic slip events beneath Nicoya Peninsula, Costa Rica, *Earth Planet. Sci. Lett.*, 410, 117-127,  
444 doi:10.1016/j.epsl.2014.11.015, 2015.

445 Fagereng, Å.: Wedge geometry, mechanical strength, and interseismic coupling of the Hikurangi subduction thrust, New  
446 Zealand, *Tectonophysics*, 507(1-4), 26-30, doi:10.1016/j.tecto.2011.05.004, 2011.

447 Fagereng, Å. and Biggs, J.: New perspectives on ‘geological strain rates’ calculated from both naturally deformed and actively  
448 deforming rocks, *J. Struct. Geol.*, 125, 100-110, doi:10.1016/j.jsg.2018.10.004, 2019.

449 Früh-Green, G. L., Orcutt, B. N., Green, S. L., Cotterill, C., and the Expedition 357 Scientists: Expedition 357 summary, *Proc.*  
450 *IODP*, 357, 1-34, doi:10.14379/iodp.proc.357.101.2017, 2017.

451 Galloway, W. E.: Depositional evolution of the Gulf of Mexico sedimentary basin, *Sediment. Basins World*, 5, 505-549,  
452 doi:10.1016/S1874-5997(08)00015-4, 2008.

453 Goldfinger, C., Ikeda, Y., Yeats, R. S. and Ren, J.: Superquakes and supercycles, *Seismolog. Res. Lett.*, 84(1), 24-32,  
454 doi:10.1785/0220110135, 2013.

455 Graham, S. E., Loveless, J. P., and Meade, B. J.: A Global Set of Subduction Zone Earthquake Scenarios and Recurrence  
456 Intervals Inferred From Geodetically Constrained Block Models of Interseismic Coupling Distributions, *Geochem. Geophys.*  
457 *Geosyst.*, 22(11), e2021GC009802, doi:10.1029/2021GC009802, 2021.

458 Hampton, M. J., Lee, H. J., and Locat, J.: Submarine landslides, *Rev. Geophys.*, 34(1), 33-59, doi:10.1029/95RG03287, 1996.

459 Hayes, G. P., Moore, G. L., Portner, D. E., Hearne, M., Flamme, H., Furtney, M., and Smoczyk, G. M.: Slab2, a comprehensive  
460 subduction zone geometry model, *Science*, 362, 58-61, doi:10.1126/science.aat4723, 2018.



461 Huang, Y.: Earthquake rupture in fault zones with along-strike material heterogeneity, *JGR Solid Earth*, 123(11), 9884-9898,  
462 doi:10.1029/2018JB016354, 2018.

463 Ito, Y., Hino, R., Kido, M., Fujimoto, H., Osada, Y., Inazu, D., Ohta, Y., Iinuma, T., Ohzono, M., Miura, S. and Mishina, M.:  
464 Episodic slow slip events in the Japan subduction zone before the 2011 Tohoku-Oki earthquake, *Tectonophys.*, 600, 14-26,  
465 doi:10.1016/j.tecto.2012.08.0222013, 2013.

466 Koppers, A. A. P., and Coggon, R. (Eds.): *Exploring Earth by Scientific Ocean Drilling: 2050 Science Framework*,  
467 *International Ocean Discovery Program*, College Station, TX, doi:10.6075/J0W66J9H, 2020.

468 Martin, R. G.: Northern and eastern Gulf of Mexico continental margin: stratigraphic and structural framework, in: *AAPG*  
469 *Studies in Geology 7: Framework, Facies, and Oil-Trapping Characteristics of the Upper Continental Margin*, edited by:  
470 Coleman, J., Bouma, A. H., and Moore, G., American Association of Petroleum Geologists, Tulsa, OK, 21-42, 1978.

471 McGuire, J.: Seismic cycles and earthquake predictability on East Pacific Rise transform faults, *Bull. Seismol. Soc. Am.*, 98,  
472 1067-1084, doi:10.1785/0120070154, 2008.

473 McNeill, L. C., Shillington, D. J., Carter, G. D. O., and the Expedition 381 Participants: Corinth Active Rift Development,  
474 *Proc. IODP*, 381, doi:10.14379/iodp.proc.381.2019, 2019.

475 Métois, M., Socquet, A., and Vigny, C.: Interseismic coupling, segmentation and mechanical behavior of the central Chile  
476 subduction zone, *J. Geophys. Res. Solid Earth*, 117(B3), B03406, doi:10.1029/2011JB008736, 2012.

477 Minson, S. E., Brooks, B. A., Glennie, C. L., Murray, J. R., Langbein, J. O., Owen, S. E., Heaton, T. H., Iannucci, R. A., and  
478 Hauser, D. L.: Crowdsourced earthquake early warning, *Sci. Adv.*, 1(3), e1500036, doi:10.1126/sciadv.1500036, 2015.

479 Moreno, M., Haberland, C., Oncken, O., Rietbrock, A., Angiboust, S., and Heidbach, O.: Locking of the Chile subduction  
480 zone controlled by fluid pressure before the 2010 earthquake, *Nature Geosci.*, 7(4), 292-296, doi:10.1038/ngeo2102, 2014.

481 Morgan, J., Gulick, S., Mellett, C. L., Green, S. L., and the Expedition 364 Scientists: Chicxulub: Drilling the K-Pg impact  
482 crater, *Proc. IODP*, 364, doi:10.14379/iodp.proc.364.2017, 2017.

483 Morris, J. T., and Renken, K. A.: Past, present, and future nuisance flooding on the Charleston peninsula, *PLOS One*, 15(9),  
484 e0238770, doi:10.1371/journal.pone.0238770, 2020.

485 Mountain, G., Proust, J.-N., McInroy, D., Cotterill, C., and the Expedition 313 Scientists: New Jersey Shallow Shelf, *Proc.*  
486 *IODP*, 313, doi:10.2204/iodp.proc.313.2010, 2010.

487 National Geophysical Data Center / World Data Service (NGDC/WDS): NCEI/WDS Global Significant Earthquake Database.  
488 NOAA National Centers for Environmental Information. doi:10.7289/V5TD9V7K, 2023a.

489 National Geophysical Data Center / World Data Service (NGDC/WDS): NGDC/WDS Global Historical Tsunami Database.  
490 NOAA National Centers for Environmental Information, doi: 10.7289/V5PN93H7, 2023b.

491 National Ocean Service (NOS) / Center for Operational Oceanographic Products & Services (CO-OPS): Tsunami Capable  
492 Tide Stations, National Oceanic and Atmospheric Administration (NOAA), <https://tidesandcurrents.noaa.gov/tsunami/>, 2023.

493 Nixon, C. W., McNeill, L. C., Bull, J. M., Bell, R. E., Gawthorpe, R. L., Henstock, T. L., Christodoulou, D., Ford, M., Taylor,  
494 B., Sakellariou, D., Ferentinos, G., Papatheodorou, G., Leeder, M. R., Collier, R. E. L. I., Goodliffe, A. M., Sachpazi, M., and

495 Kranis, H.: Rapid spatiotemporal variations in rift structure during development of the Corinth Rift, central Greece, *Tectonics*,  
496 35(5), 1225-1248, doi:10.1002/2015TC004026, 2016.

497 Özel, N. M., Necmioglu, O., Ergintav, S., Özel, O., Italiano, F. et al.: MARSite-Marmara Supersite: accomplishments and  
498 outlook, *EGU General Assembly 2017*, 23-28 April 2017, Vienna, Austria, EGU2017-18891, 2017.

499 Pan, S., Naliboff, J., Bell, R., and Jackson, C.: Bridging spatiotemporal scales of normal fault growth during continental  
500 extension using high-resolution 3D numerical models, *Geochem. Geophys. Geosyst.*, 23, e2021GC010316,  
501 doi:10.1029/2021GC010316, 2022.

502 Popenoe, P., Schmuck, E. A., and Dillon, W. P.: The Cape Fear landslide: slope failure associated with salt diapirism and gas  
503 hydrate decomposition, in: *Submarine Landslides: Selected Studies in the U.S. Exclusive Economic Zone*, edited by: Schwab,  
504 W. C., Lee, H. J., and Twichell, D. C., U.S. Geological Survey, Washington, D.C., 40-53, 1993.

505 Ruff, L. J.: Do trench sediments affect great earthquake occurrence in subduction zones? *Pure Appl. Geophys.*, 129(1-2), 263-  
506 282, doi:10.1007/BF00874629, 1989.

507 Ruiz, S., Metois, M., Fuenzalida, A., Ruiz, J., Leyton, F., Grandin, R., Vigny, C., Madariaga, R. and Campos, J.: Intense  
508 foreshocks and a slow slip event preceded the 2014 Iquique M w 8.1 earthquake, *Science*, 345(6201), 1165-1169,  
509 doi:10.1126/science.1256074, 2014.

510 Sainsbury, E. M., Schiemann, R. K. H., Hodges, K. I., Shaffrey, L. C., Baker, A. J., and Bhatia, K. T.: How important are post-  
511 tropical cyclones for European windstorm risk? *Geophys. Res. Lett.*, 47(18), e2020GL089853, doi:10.1029/2020GL089853,  
512 2020.

513 Saito, T., and Noda, A.: Mechanically coupled areas on the plate interface in the Nankai Trough, Japan and a possible seismic  
514 and aseismic rupture scenario for megathrust earthquakes, *JGR Solid Earth*, 127(8), e2022JB023992,  
515 doi:10.1029/2022JB023992, 2022.

516 Sawyer, D. E., and DeVore, J. R.: Elevated shear strength of sediments on active margins: evidence for seismic strengthening,  
517 *Geophys. Res. Lett.*, 42(23), 10216-10221, doi:10.1002/2015GL066603, 2015.

518 Scharroo, R., Smith, W. H. F., and Lillibridge, J. L.: Satellite altimetry and the intensification of Hurricane Katrina, *Eos Trans.*  
519 *AGU*, 86(40), 366, doi:10.1029/2005EO400004, 2005.

520 Scholz, C. H.: *The Mechanics of Earthquakes and Faulting* (3<sup>rd</sup> ed.), Cambridge University Press, Cambridge U.K., 2019.

521 Schulten, I., Mosher, D. C., Piper, D. J. W., and Krastel, S.: A Massive Slump on the St. Pierre Slope, A New Perspective on  
522 the 1929 Grand Banks Submarine Landslide, *JGR Solid Earth*, 124(8), 7538-7561, doi:10.1029/2018JB017066, 2019.

523 Sparkes, R., Tilmann, F., Hovius, N., and Hillier, J.: Subducted seafloor relief stops rupture in South American great  
524 earthquakes: Implications for rupture behaviour in the 2010 Maule, Chile earthquake, *Earth Planet. Sci. Lett.*, 298(1-2), 89-  
525 94, doi:10.1016/j.epsl.2010.07.029, 2010.

526 Stein, S., Geller, R. J., and Liu, M.: Why earthquake hazard maps often fail and what to do about it, *Tectonophys*, 562-563, -  
527 25, doi:10.1016/j.tecto.2012.06.047, 2012.

528 Taylor, S. K., Bull, J. M., Lamarche, G., and Barnes, P. M.: Normal fault growth and linkage in the Whakatane Graben, New  
529 Zealand, during the last 1.3 Myr, *J. Geophys. Res.*, 109, B02408, doi:10.1029/2003JB002412, 2004.

530 Telesca, L.: Time-clustering of natural hazards, *Nat. Hazards*, 40, 593-601, doi:10.1007/s11069-006-9023-z, 2007.

531 Urlaub, M., Talling, P. J., and Masson, D. G.: Timing and frequency of large submarine landslides: implications for  
532 understanding triggers and future geohazard, *Quat. Sci. Rev.*, 72, 63-82, doi:10.1016/j.quascirev.2013.04.020, 2013.

533 van Rijnsingen, E., Lallemand, S., Peyret, M., Arcay, D., Heuret, A., Funicello, F., and Corbi, F. (2018). How subduction  
534 interface roughness influences the occurrence of large interplate earthquakes, *Geochem. Geophys. Geosyst.*, 19, 2342-2370,  
535 doi:10.1029/2018GC007618, 2018.

536 Vega, A. J., Miller, P. W., Rohli, R. V., and Heavilin, J.: Synoptic climatology of nuisance flooding along the Atlantic and  
537 Gulf of Mexico coasts, USA, *Nat. Hazards*, 105, 1281-1297, doi:10.1007/s11069-020-04354-5, 2021.

538 Wallace, L. M., Barnes, P., Beavan, J., Van Dissen, R., Litchfield, N., Mountjoy, J., Langridge, R., Lamarche, G., and Pondard,  
539 N.: The kinematics of a transition from subduction to strike-slip: An example from the central New Zealand Plate boundary,  
540 *JGR Solid Earth*, 117(B2), B02405, doi:10.1027/2011JB008640, 2012.

541 Wallace, L. M., Saffer, D. M., Barnes, P. M., Pecher, I. A., Petronotis, K. E., LeVay, L. J., and the Expedition 372/375  
542 Scientists: Hikurangi Subduction Margin Coring, Logging, and Observatories, *Proc. IODP*, 372B/375,  
543 doi:10.14379/iodp.proc.372B375.2019, 2019.

544 Walsh, K. J. E., McBride, J. L., Klotzbach, P. J., Balachandran, S., Camargo, S. J., Holland, G., Knutson, T. R., Kossin, J. P.,  
545 Lee, T., Sobel, A., and Sugi, M.: Topical cyclones and climate change, *WIREs Climate Change*, 7(1), 65-89,  
546 doi:10.1002/wcc.371, 2016.

547 Wang, K., and Bilek, S. L.: Do subducting seamounts generate or stop large earthquakes?, *Geology*, 39(9), 819-822,  
548 doi:10.1130/G31856.1, 2011.

549 Webster, J.M., Yokoyama, Y., Cotterill, C., and the Expedition 325 Scientists: Great Barrier Reef Environmental Changes,  
550 *Proc. IODP*, 325, doi:10.2204/iodp.proc.325.2011, 2011.

551 Wells, D. L., and Coppersmith, K. J.: New empirical relationships among magnitude, rupture length, rupture width, rupture  
552 area, and surface displacement, *B. Seismol. Soc. Am.*, 84(4), 974-1002, doi:10.1785/BSSA0840040974, 1994.