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Attribution of 2022 early-spring heatwave in India and Pakistan to climate change: lessons in assessing vulnerability and preparedness in reducing impacts

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1 Attribution of 2022 early-spring heatwave in India and Pakistan to
 2 climate change: Lessons in assessing vulnerability and preparedness
 3 in reducing impacts.

4
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46 Abstract

47 In March 2022, large parts over the north Indian plains including the breadbasket region, and southern
48 Pakistan began experiencing prolonged heat, which continued into May. The event was exacerbated
49 due to prevailing dry conditions in the region, resulting in devastating consequences for public health
50 and agriculture. Using event attribution methods, we analyse the role of human-induced climate change
51 in altering the chances of such an event. To capture the extent of the impacts, we choose March-April
52 average of daily maximum temperature over the most affected region in India and Pakistan as the
53 variable. In observations, the 2022 event has a return period of ~1-in-100 years. For each of the climate
54 models, we then calculate the change in probability and intensity of a 1-in-100 year event between the
55 actual and counterfactual worlds for quantifying the role of climate change. We estimate that human-
56 caused climate change made this heatwave about 1°C hotter and 30 times more likely in the current,
57 2022 climate, as compared to the 1.2 °C cooler, pre-industrial climate. Under a future global warming
58 of 2°C above pre-industrial levels, heatwaves like this are expected to become even more common (2–
59 20 times more likely) and hotter (by 0-1.5°C) compared to now. Stronger and frequent heat waves in
60 the future will impact vulnerable groups as conditions in some regions exceed limits for human
61 survivability. Therefore, mitigation is essential for avoiding loss of lives and livelihood. Heat Action
62 Plans (HAPs) have proved effective to help reduce heat-related mortality in both countries.

63

64 Keywords

65 Extreme event attribution, heatwave, climate change, India, Pakistan

66

67 1. Introduction

68

69 Towards the end of spring of 2022, large parts of South Asia including India and
70 Pakistan began experiencing prolonged periods of hot weather, which continued into summer.
71 The month of March was the hottest in India since observed records began in 1901, according
72 to the India Meteorological Department (IMD; Madaan, 2022). Temperatures were
73 consistently 3°C–8°C above the long period average (L.P.A, 1981–2010 climatology),
74 breaking many decadal and some all-time records in several parts of the country, including the
75 western Himalayas, the plains of Punjab, Haryana, Delhi, Rajasthan and Uttar Pradesh (IMD,
76 2022). The states of Odisha, Madhya Pradesh, Gujarat, Chhattisgarh, Telangana and Jharkhand
77 also experienced heatwaves, in some areas quite severe, with temperatures ranging from 40°C–
78 44°C in the last days of March (IMD, 2022). In Pakistan many individual weather stations
79 recorded monthly all-time highs in March (Pakistan Meteorological Department (PMD),
80 2022b). The heatwave conditions persisted into April, reaching a preliminary peak towards the
81 end of the month. Around 300 large forest fires occurred in India on April 28, a third of these
82 in the state of Uttarakhand (Rajeevan, 2023). These months were also extremely dry, with
83 rainfall at 62% and 73.6% below the L.P.A over Pakistan (PMD, 2022a), and 71% and 3%
84 below L.P.A over India for March and April, respectively (IMD, 2022b, 2022a). By April 29,
85 almost 70% of India was affected by the heatwave. Temperatures above 49°C were recorded
86 in Jacobabad in Sindh, Pakistan (Ilyas, 2022), and 30% of the country was gripped by the
87 heatwave. Towards the end of April and in May, the heatwave extended into the coastal areas
88 and eastern parts of India (IMD, 2022).

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Even though inhabitants in these regions are used to high ambient temperatures, temperatures exceeding 40°C can increase mortality (Desai et al., 2015; Rathi & Sodani, 2021; Rathi et al., 2021). The heatwave coincided with the holy Ramadan period, which affected the coping capacities of those fasting, particularly in Pakistan, thus exacerbating health impacts. The updated death toll for the 2022 from India event is 33 (Datt, 2023) which is 8 more than the initial estimate. The number of heat-related deaths recorded in Pakistan, where temperature records were broken in several places remain at 65 (Irfan, 2023). A particularly notable effect is the impact on wheat crops and yields in the wheat-growing regions in the northern plains of India and Southern Pakistan. Anomalously high temperatures during the wheat harvest season in these parts (February-May) are known to adversely affect grain filling and cause early senescence (Lobell et al., 2012), thereby reducing yields (Zachariah et al., 2021). The country had been aiming for 111.32 million tonnes of wheat for 2022-23; however, the actual production was 106.84 million tonnes, with the shortfall of ~20% (Arora & Bhardwaj, 2022; Gupta, 2023). The export ban imposed by India on wheat due to concerns about domestic food security reportedly added further stress on global food prices and food security in an already tight market given the war in Ukraine (AFP, 2022). At least 16 glacial lake outbursts in Pakistan during 2022 was linked to the heatwave as compared to annual average of 5-6 (Fox, 2022; Janjua, 2022). Fig. 1 illustrates the far-reaching impacts of the 2022 extreme heat event, and in particular the varied ways in which the heightened (lessened) vulnerability and exposure to the event is expected to have heightened (lessened) the associated impacts.

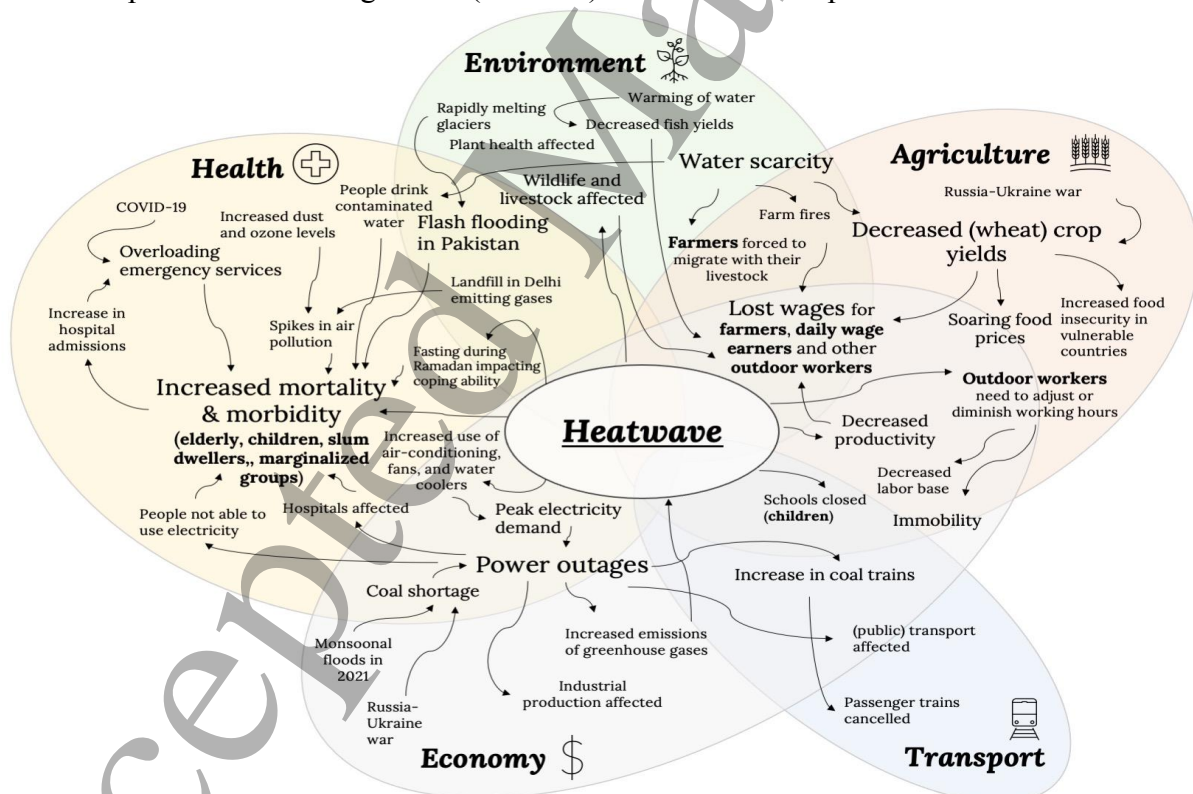


Fig. 1: Conceptual map of impact pathways during the heatwave.

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115 While heatwaves are not uncommon in this part of the world during the later months of
116 the pre-monsoon (MAMJ) season (Chaudhury et al., 2000; Sharma & Mujumdar, 2017; Zahid
117 & Rasul, 2012; IMD Climate Summary, 2015; 2016), an increase in such conditions in the
118 earlier, March-April period is observed in recent decades (Singh et al., 2021; Zahid & Rasul,
119 2012). The heatwave during March–April 2022 was characterised by anomalously high
120 temperatures in large parts of India and Pakistan (Fig. 2(a)) due to the persistence of an upper
121 atmospheric high-pressure system (anticyclone) (National Weather Forecasting Centre, 2022a,
122 2022b). In a recent study Rashid et al. (2022) found that in early summer, La Niña conditions
123 favour development of anticyclonic systems in this region. This may have been the case during
124 the 2022 event considering that the Pacific was in the La Niña phase at that time.

125 Another important feature of this event was the extremely dry conditions that
126 accompanied the hot weather, thereby making the conditions favourable for enhanced surface
127 heating. Upper-level synoptic scale systems embedded in the subtropical westerly jet stream
128 called western disturbances (Hunt et al., 2018, 2019) that are responsible for precipitation
129 during December-April in northwest India and Pakistan were absent during March and April,
130 2022, causing a large rainfall deficit during this period (Express News Service, 2022; Fig.
131 2(b)). Thus, the 2022 heatwave was a dry event, in contrast to the previous, major heatwave of
132 2015 that occurred later in the season in June, concurring with high humidity levels and
133 resulting in 3500 direct heat-related deaths in both countries (Saeed, et al., 2021).

134
135 Important concerns raised in the immediate aftermath of the event include (i) whether
136 and to what extent can the 2022 heatwave be attributed to climate change (ii) whether
137 heatwaves such as this one will become worse in the future, and (iii) what anticipatory
138 strategies can be implemented for dealing with such extremes in the future. In this study, we
139 attempt to answer these questions systematically, using the peer-reviewed protocol for event
140 attribution (EA) and vulnerability and exposure (V&E) analyses developed by Philip et al.,
141 (2020). In keeping with the unusual timing of the heat episode, which occurred earlier in the
142 year, the spatial extent and the associated widespread impacts, we define the event as the
143 March-April average of daily maximum temperature, over the north Indian plains to the west
144 of the Himalayas and the lowlands in southern Pakistan to the east of the Sulaiman range
145 (highlighted in blue in Fig. 2). We choose to focus on the maximum air temperature, as opposed
146 to a more complex metric for heat stress because (i) historical observations of humidity are less
147 reliable (and are less available) than temperature measurements in the relevant regions, (ii) the
148 question of model fidelity becomes more complex when assessing multivariate extreme events
149 (Cannon et al., 2020; Sippel et al., 2016), and (iii) the concurrence of positive anomalies in
150 temperature with negative rainfall anomalies (Fig. 2(a-b)) during March-April 2022 over India
151 and Pakistan suggests that the event exhibits the characteristics of a “dry heatwave”.

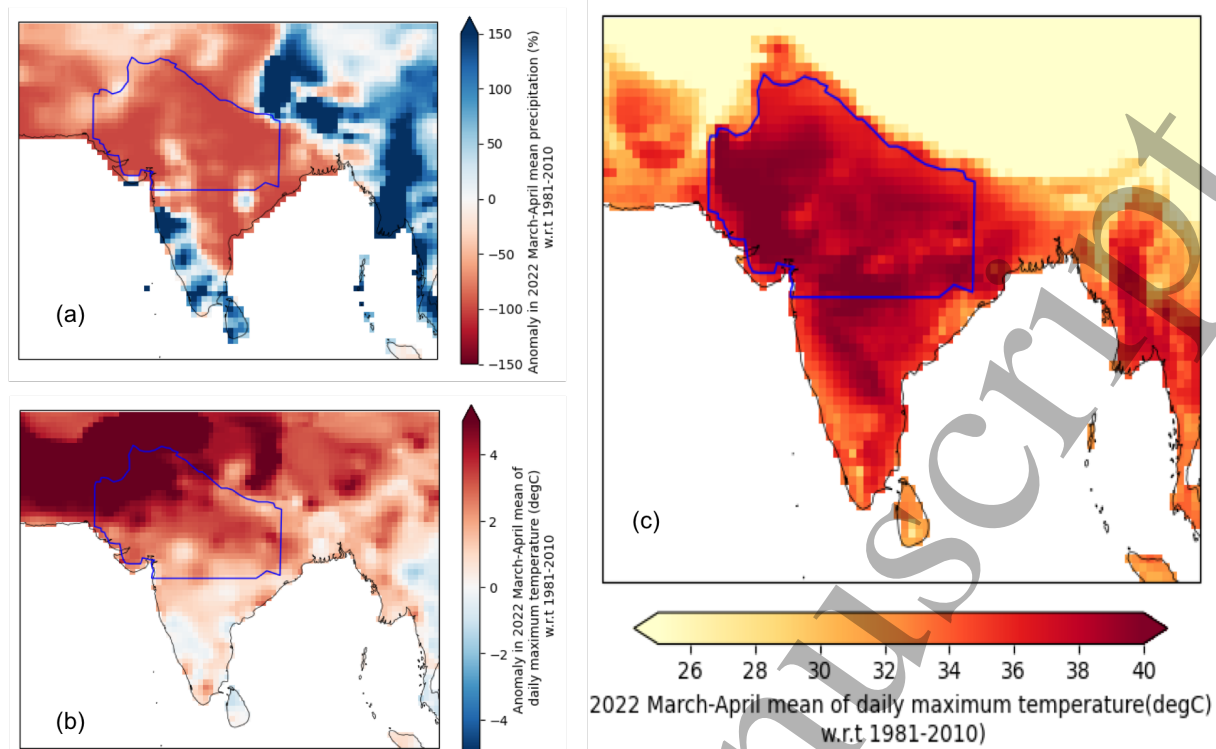


Fig. 2: (a) Percentage deviation of observed (CPC) precipitation during Mar-Apr 2022 from the 1981-2010 climatology (%). (b) Observed (CPC) anomaly in March-April average daily maximum temperature for the year 2022 w.r.t. 1979-2022 climatology ($^{\circ}\text{C}$). The study region is highlighted by the blue polygon. (c) March-April average of observed (CPC) maximum temperature for the year 2022 ($^{\circ}\text{C}$).

Over South Asia and other global regions, there is *high confidence* that the likelihood and intensity of extreme heat has significantly increased (Chakraborty et al., 2018; Dimri, 2019; Donat et al., 2016; Dunn et al., 2020; Rohini et al., 2016; Roy, 2019; Seneviratne et al., 2021; Sheikh et al., 2015; Zahid & Rasul, 2012) and there is *robust evidence* confirming the role of human-induced climate change in driving them (Dileepkumar et al., 2018; Dileepkumar et al., 2021; Pattanayak et al., 2017; Seong et al., 2021; van Oldenborgh et al., 2018; Wang et al., 2017; Wehner et al., 2016). Nonetheless, at smaller scales, local factors can further exacerbate or alleviate the event characteristics. In parts of India, warming signals are attenuated due to the increased concentration of reflective sulfate aerosols in the atmosphere from burning fossil fuels, and the cooling from agricultural intensification and high irrigation activity in the region (Mishra et al., 2020; Thiery et al., 2017, 2020; van Oldenborgh et al., 2018). On the other hand, absorbing aerosols, particularly Black Carbon (BC) that also results from fuel combustion and crop-burning, is found to intensify high temperatures in these areas (Mondal et al., 2021). However, it should be noted that the studies that examine the link between irrigation and regional cooling account for soil moisture at field capacity or as a percentage of soil saturation, and use annual irrigated areas, thereby overlooking the fact that pre-monsoonal irrigation activities in India are only minimal when compared to the major monsoon (Kharif) and winter (Rabi) cropping seasons (Jha et al., 2022; Devanand et al., 2019). Combined, this evidence suggests that for the specific case of a March–April heatwave over these parts, the importance of increasing irrigation in suppressing the warming effect of greenhouse gases might be smaller than previously thought.

180 2. Data and Models

181

182 *2.1 Observational data*

183 Gridded datasets for daily maximum temperature provided by the National Oceanic and
184 Atmospheric Administration (NOAA) Climate Prediction Centre (CPC), available at $0.5^\circ \times$
185 0.5° resolution for the period 1979–present (from
186 <https://psl.noaa.gov/data/gridded/data.cpc.globaltemp.html>) are used as the primary dataset for
187 observational analysis. Additionally, we use gridded datasets of observed daily maximum
188 temperature at $1^\circ \times 1^\circ$ resolution for the period from January 1, 1951 to April 30, 2022 provided
189 by the India Meteorological Department (IMD; available at
190 https://www.imdpune.gov.in/Clim_Pred_LRF_New/Gridded_Data_Download.html) as a
191 supplementary observational product although its spatial extent is limited to within the
192 geographical borders of India. Both datasets are interpolated from station data using Shepard's
193 interpolation algorithm (Chen et al., 2008; Pai et al., 2014; Srivastava et al., 2009; Xie et al.,
194 2007).

195

196 For studying the effect of climate change on temperature, we assume that the location
197 parameter of the best-fitted probability distribution for temperature varies with the Global
198 Mean Surface Temperature (GMST), an accepted measure of anthropogenic climate change
199 (e.g., Luu et al., 2021; van Oldenborgh et al., 2017). To this end, we use low-pass filtered
200 estimates of GMST from the National Aeronautics and Space Administration (NASA)
201 Goddard Institute for Space Science (GISS) surface temperature analysis (GISTEMP, Hansen
202 et al., 2010; Lenssen et al., 2019) as the covariate.

203

204 *2.2. Reliability of the observed datasets for the study region*

205

206 Fig. 3(a-b) shows the linear trends in March–April average daily maximum temperature for the
207 period 1979–2022, from the CPC and IMD datasets, respectively. We find strong positive
208 trends in this season almost everywhere in the study domain. The CPC dataset exhibits negative
209 trends further south (Fig. 3(a)) as opposed to no significant trends in the IMD data (Fig. 3(b)).
210 This suggests the need to consider both datasets in the analysis, to account for the uncertainties
211 from differences in the trends. Fig. 3(c-d) shows the time series of March–April average daily
212 maximum temperature, averaged over the study region, from CPC (1979-present; Fig. 3(c))
213 and for the Indian part of the region only from IMD (1951-present; Fig. 3(d)). Overall, these
214 datasets agree with each other in terms of magnitudes, year-to-year variability and the positive
215 trend between 1979-2022, thus attesting to the homogeneity in the larger region, and further
216 justifying their use as complementing datasets for the rest of the analysis.

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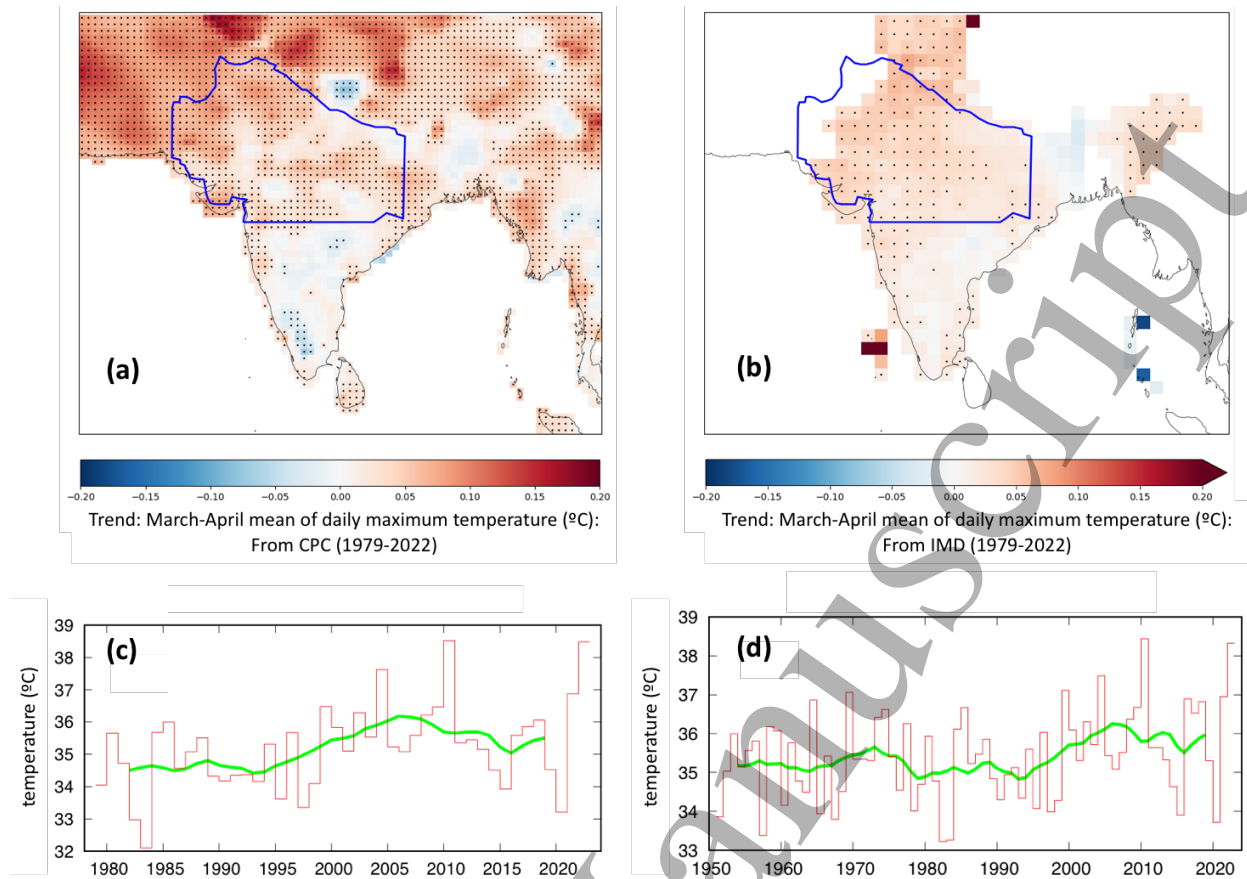


Fig. 3: Trends in observed daily maximum temperature averaged over March-April from the (a) CPC and (b) IMD datasets, considering the years 1979-2022. Stippling indicates trends that are significant at the 10% significance level. Time series of area-averaged March-April average of observed daily maximum temperature along with the ten-year running mean (shown by the green line) based on the (c) CPC and (d) IMD datasets.

2.3 Model simulations

We use multi-model ensembles from six climate modelling projects that use different framings such as Sea Surface temperature (SST) driven global circulation high resolution models, coupled global circulation models and regional climate models. Specifically, we consider simulations from the following projects: (i) Half a degree Additional warming Prognosis and Projected Impacts (HAPPI; Wehner et al., 2018), (ii) Coordinated Regional Climate Downscaling Experiment (CORDEX-CORE; Teichmann et al., 2021), (iii) IPSL-CM6A-LR global climate model large ensemble (Bonnet et al., 2021; Boucher et al., 2020), (iv) GFDL Forecast-oriented Low Ocean Resolution version of CM2.5 (FLOR; Vecchi et al., 2014), (v) High Resolution Model Intercomparison Project (HighResMIP; Haarsma et al., 2016) and (vi) Coupled Model Intercomparison Project version 6 (CMIP6; Eyring et al., 2016). More details about these models and the experiments that are used in this study are provided in Supplementary Section S1.

The 1979–2022 period for which the observed data (CPC) is available is chosen for model evaluation, while the entire length of simulations up to the year 2022 is considered for the attribution analysis. As with observations, for the SST-forced simulations, we use observed

242 GMST as covariate, whereas for the coupled models, we use the corresponding model-based
243 GMST estimates.

245 3. Methods

246
247 We use a standard, peer-reviewed probability-based framework developed for rapid attribution
248 assessments (Philip et al. 2020), for quantifying whether and to what extent the frequency
249 and/or magnitude of a class of extremes at least as extreme as the event of interest is attributable
250 to climate change. Such results can help inform decision-makers in planning
251 adaptation/mitigation strategies against future impacts. This is complementary to the ‘storyline
252 approach’ which is also interesting in itself, separating the climate change contributions to the
253 thermodynamic and the dynamic aspects of the event (Otto et al., 2016; Vautard et al., 2016).
254 Although the storyline approach is important from a research perspective, this is not easily
255 automated and therefore, not within the immediate scope of a rapid attribution study.

256
257 For quantifying attribution of the event to climate change, we calculate its return period and
258 the changes in the event's probability (given by the probability ratio PR) and intensity between
259 the climate of today ($clim_1$) and a hypothetical past before anthropogenic activities began
260 altering the climate ($clim_0$). PR is the ratio of the probability of an event as strong or stronger
261 than the event of interest in $clim_1$ to its probability in $clim_0$ and can take values in $(0, \infty)$. A
262 value of PR more (less) than 1 implies that the event is made more (less) likely by human-
263 induced climate change whereas $PR = 1$ suggests that there is no evidence of climate change
264 in the likelihood of the event. Following the same framing, the change in intensity is the
265 difference in the variable threshold for a given probability or return period between $clim_1$ and
266 $clim_0$. See Supp. Fig. S1 for an illustration of how these metrics are calculated.

267
268 In the main method, we fit the data with a non-stationary probability distribution with GMST
269 as covariate. This distribution is then shifted up to the 2022 climate when the event was
270 observed and shifted down to the pre-industrial (late 19th century) climate that is 1.2°C below
271 the 2022 levels (see Global Warming Index; <https://www.globalwarmingindex.org>).
272 Additionally, we analyse the expected changes in probability and intensity in a future warmer
273 climate scenario that is +2.0°C above pre-industrial GMST levels (or 0.8°C warmer than 2022
274 levels).

275
276 We select the Gaussian distribution for modelling the March-April mean of daily maximum
277 temperature in the study region. This is a justified choice, on account of the short length of the
278 CPC dataset resulting in sparsely populated tails (Philip et al., 2018). The Gaussian distribution
279 is assumed to shift due to global warming without changing its shape. This is a first-order
280 assumption that is made when performing attribution analyses of temperature events. The
281 assumption has been validated in past studies, using climate models with long length of data
282 by checking the past and present distributions from the non-stationary fits against distributions
283 of the past and present climate in independent time slices (e.g., Kew et al. 2019; van
284 Oldenborgh et al. 2018; Uhe et al. 2017). The shift is factored into the analysis by linearly

285 varying the location parameter (μ) of the distribution with (low-pass filtered) global mean
286 surface temperature while holding the scale parameter (σ) constant, as shown in Eq. (1).

287

$$\mu = \mu_o + \alpha T \text{ and } \sigma = \sigma_o \quad (1)$$

288

289 where T is the 4-year smoothed GMST anomaly, and α is the trend.

290 This framework is detailed in the peer-reviewed World Weather Attribution protocol
291 ([https://www.worldweatherattribution.org/pathways-and-pitfalls-in-extreme-event-](https://www.worldweatherattribution.org/pathways-and-pitfalls-in-extreme-event-attribution)
292 [attribution](https://www.worldweatherattribution.org/pathways-and-pitfalls-in-extreme-event-attribution); Philip et al., 2020; van Oldenborgh et al., 2021) and has been successfully applied
293 in recent studies (Ciavarella et al., 2021; Luu et al., 2021).

294

295 In addition to the above method, we use simulations from three fixed climate model
296 experiments from the HAPPI project- one for current conditions (Hist), and two counterfactuals
297 for the world that would have been in the absence of anthropogenic emissions (Nat) and a
298 +2.0°C warmer world with adjusted concentrations of CO₂, other greenhouse gases, and
299 aerosols. For the observed return period (or probability) of the event in the 2022 climate, we
300 first obtain the magnitude in the Hist climate for each of the models. Thereafter, the return
301 period of this magnitude in the counterfactual scenario is calculated for the respective models.
302 The PR is the ratio of the probability (or return period) of the event in +2.0°C warmer world to
303 that in Hist. The intensity change attributable to climate change is the difference between these
304 event magnitudes.

305 Finally, results from observations and the models that pass the validation tests are synthesised
306 into a single attribution statement. See Philip et al., (2020), Ciavarella et al., (2021) and Li &
307 Otto, (2022) for details.

308

309 4. Results and Discussions

310 4.1 Observational analysis: trend and return period

311 The left panels in Fig. 4 show the response of area-averaged March-April mean daily maximum
312 temperatures to the global mean surface temperature anomaly, for the CPC and IMD datasets.
313 Despite the limited length of the data series, there are clear trends in temperature for both
314 datasets (Fig. 4(a,c)) that suggest warming of the late spring and early summer in the study
315 region consistent with global warming. Fitting these time series to a Gaussian distribution
316 shifting with GMST (Section 3) allows us to calculate the return period of the 2022 temperature
317 in the 2022 climate as well as a past, 1.2°C cooler climate, along with the change in intensity
318 of the event between these climates for a quantitative appraisal of the role of climate change in
319 causing this event.

320

321 The right panels (Fig. 4(b,d)) show the return period curves in the present, 2022 climate and
322 the past climate when the global mean temperature was 1.2°C cooler, for the two datasets,
323 along with their confidence intervals (CI). The two-sided 95% CI for these curves is estimated

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3 324 using a non-parametric bootstrap, by repeating the fit 1000 times with (variable, covariate)
4 325 pairs drawn from the original series, with replacement. Although the CPC dataset has high
5 326 resolution over the study region, it is too short to estimate the return period with confidence.
6 327 Therefore, we compare these estimates with those based on IMD dataset for the part of the
7 328 study region in India. Upon fitting Gaussian distributions, the best estimates of the return period
8 329 of the 2022 event in the current climate emerges as 1-in-130 year (uncertainty: 1-in- 8000 to 1-
9 330 in-20 year) for the study region based on CPC (Fig. 4(b)), and 1-in-103 year (uncertainty: 1-in-
10 331 1250 to 1-in-25 year) for the region over India based on IMD (Fig. 4(d)) and 1-in-110 year
11 332 (uncertainty: 1-in- 4500 to 1-in-15 year) for the same region, again , based on CPC (Table S5).
12 333 These numbers imply that, based on available observed data, the March-April average
13 334 maximum temperature in the study region during 2022 is best defined as a 1-in-100 year event
14 335 in the current climate.
15 336

16 337 This event would have been highly unlikely in a world without climate change, with very high
17 338 return periods of 20,000 and 1,500 years with wide uncertainty bounds (see Supplementary
18 339 Table S5), in the CPC and IMD datasets, respectively, indicating that climate change has
19 340 increased the chances of the event. However, these numbers are many times higher than the
20 341 length of the observational data (44 years long for CPC and 73 for IMD) and the 2022 value
21 342 is an outlier in both the 1.2°C cooler and 2022 climates. Therefore, we cannot rely on the
22 343 precision of these numbers. A possible reason for these high numbers and the large
23 344 uncertainties in the return period of the 2022 event could be due to the Gaussian distribution
24 345 not being able to model the tail of the data distribution as seen in Fig. 4(b,d). The large
25 346 uncertainties are also reflected in the PR and ΔI estimates from the datasets. The best estimate
26 347 for PR between the 2022 climate and the 1.2°C cooler climate is 155 (0.9 to 1,000,000 with
27 348 95% CI; Table 1) and 15 (0.7 to 600; Table 1), from the CPC and IMD datasets, respectively,
28 349 with change in intensity (ΔI) estimates of 1.7°C (-0.019°C ... 3.7°C) and 0.92°C (-0.12°C ...
29 350 2.0°C), respectively (Table 1). These bounds encompassing no change suggests that the trend
30 351 may not yet have emerged from the noise in the observations even though the probability has
31 352 changed. This is partially due to the short length of observations resulting in larger sampling
32 353 uncertainties, along with other confounding factors including natural variability. Therefore,
33 354 we repeat this analysis climate models as well, as discussed in Section 4.2.
34 355

35 356 For testing the reliability of the Gaussian assumption, we separately fitted a Gaussian model
36 357 and a Generalized-Pareto distribution (GPD; with a threshold of 90% of the data) to simulations
37 358 of a multi ensemble climate model with long runs- the IPSL-CM6A-LR ensemble (see Section
38 359 2.2 and Supplementary Section S1 for more details about the model), Fig. S9 shows the two
39 360 fits- the Gaussian and the GPD (with a threshold of 90% of data) distributions. Although the
40 361 the best estimate of the return period shows the same order of magnitude as with the
41 362 observations (~100 years), there is a subtle departure from the data in the fitted Gaussian
42 363 distribution far tail and the fit does not capture the curvature of the data, while this is captured
43 364 by the GPD distribution. As a result, interestingly, this induces a large change in the probability
44 365 ratio, with a best estimate of about 50 in the Gaussian case and 3300 in the GPD case. The
45 366 results in terms of intensity changes are however not much changed (~1°C). This result
46 367 suggests a potential underestimation of the PRs by the Gaussian model for temperatures.

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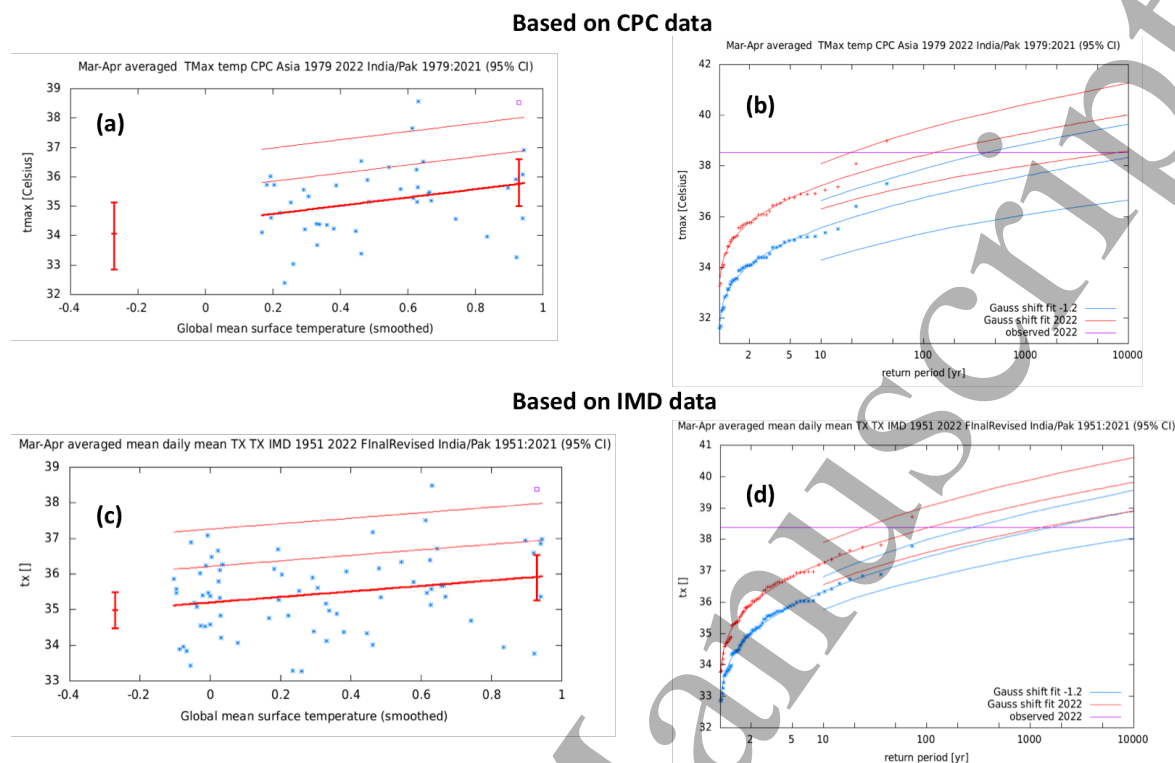


Fig. 4: (a) Response of March-April mean daily maximum temperature averaged over the study region estimated from CPC records to change in global mean temperature. The thick red line denotes the time-varying mean, and the thin red lines show 1 standard deviation (s.d) and 2 s.d above. The vertical red lines show the 95% confidence interval for the location parameter, for the current, 2022 climate and the hypothetical, 1.2°C cooler climate. The 2022 observation is highlighted with the magenta box. (b) Return periods for the 2022 climate (red lines) and the 1.2°C cooler climate (blue lines with 95% CI), based on CPC data. (c) same as (a), but using IMD data. (d) same as (b), using IMD data.

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380 4.2 Climate model analysis

381 We repeat the above analyses based on observed data with temperature simulations from the
382 participating models (discussed in Section 2.2) for estimating the model-based probability
383 ratios and intensity changes. First, we evaluate the model simulations against observations for
384 their suitability for the attribution analysis. The evaluation period considered is 1979-2022, the
385 period common to the CPC and IMD observational datasets. The climate models are evaluated against
386 the observations in their ability to capture-

387
388 1. *Seasonal cycle*: The seasonal cycle of the daily maximum temperature for the study region
389 from the climate models are qualitatively compared against the cycle from CPC dataset. A
390 model is labelled as ‘good’ if the model-based cycles capture the shape and the seasonality of
391 the observed seasonal cycle. It is labelled ‘reasonable’ if either the peaks are not well-defined
392 or if the seasonality is out of phase. If the peaks are ill-defined and the seasonality is out of
393 phase, the model is labelled as ‘bad’.

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4 395 2. *Spatial pattern*: The spatial pattern of March-April average maximum temperatures for a
5 396 larger region spanning 5°N-40°N, 60°E-100°E and encompassing the study region (Fig. 2)
6 397 from the model simulations are qualitatively compared with the pattern based on observed data.
7 398 Depending on how well the models are able to replicate the observed patterns- i.e. the north-
8 399 south gradient in temperatures and patterns due to higher temperatures in arid/semi-arid regions
9 400 (Fig. S2 (b)) - completely, at least in part or poorly, these are classified as ‘good’, ‘reasonable’
10 401 or ‘bad’, respectively.
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14 403 3. *Distribution parameters*: We check if the parameters of the fitted statistical distribution
15 404 (Gaussian shifting with GMST for this study) from the model simulations are compatible with
16 405 those from observations. A model is labelled as ‘good’ if the model parameter range lies within
17 406 the observational range (95% confidence interval), ‘reasonable’ if the ranges overlap, and
18 407 ‘bad’, if they diverge.
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22 409 A model is given an overall rating of ‘good’ if it is rated ‘good’ for all three characteristics. If
23 410 there is at least one ‘reasonable’, then its overall rating will be ‘reasonable’ and ‘bad’ if there
24 411 is at least one ‘bad’. Supplementary Table S6 summarises the model evaluation results for 66
25 412 model simulations from the various experiments. The seasonal cycles and spatial patterns for
26 413 the observed dataset and the participating models that are used for the model evaluation are
27 414 shown in Supplementary Section S2, Figs. S2-S8.

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30 415 24 out of the 66 models that have an overall rating of ‘good’ (highlighted in green in Table S5)
31 416 and are selected for the attribution analysis. Out of these, there are 22 models that have
32 417 simulations for the past 1.2°C cooler climate and 19 models that cover the future 2°C warmer
33 418 world.
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36 420 Table 1(a) (highlighted in blue) shows the PR and ΔI of the 1-in-100 year event for the present
37 421 climate, relative to a 1.2°C cooler climate, based on the observations and the 22 selected
38 422 models. Table 1(b) (highlighted in red) shows these values for a future 2°C warmer world
39 423 relative to the present climate, from the 19 models. The individual model estimates of the
40 424 change in event likelihood and event intensity are strongly correlated to each other i.e., models
41 425 showing a large increase in the intensity of 100-year hot events also have correspondingly large
42 426 probability ratios and vice versa (see supplementary Fig. S10). In the next Section, we
43 427 synthesize the results for observations and the models that pass validation.
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428 **Table 1:** (a) Probability ratio and change in intensity when compared with a 1.2°C cooler climate, for models that passed the validation tests. (b) Projected probability ratio
 429 and change in intensity when compared with a 2°C warmer climate, for models that passed the validation tests.

Model / Observations	(a) Present vs. past		(b) Future vs. present	
	Probability ratio PR	Change in intensity ΔI [°C]	Probability ratio PR	Change in intensity ΔI [°C]
CPC (1979-2022)	1.5e+2 (0.94 ... 1.2e+6)	1.7 (-0.019 ... 3.7)	-	-
IMD (1951-2022)- India only	15 (0.70 ... 6.0e+2)	0.92 (-0.12 ... 2.0)	-	-
ECEARTHr12-COSMOcrCLIM rcp85 (1)	35 (3.0 ... 1.7e+3)	1.6 (0.53 ... 2.7)	7.0 (5.0 ... 11)	1.1 (0.95 ... 1.2)
MPIr1-COSMOcrCLIM rcp85 (1)	1.1e+2 (11 ... 2.8e+3)	1.7 (0.87 ... 2.4)	11 (7.0 ... 19)	1.3 (1.2 ... 1.4)
NOESMr1-COSMOcrCLIM rcp85 (1)	14 (0.70 ... 3.5e+2)	1.2 (-0.16 ... 2.4)	6.0 (4.0 ... 9.0)	0.97 (0.84 ... 1.1)
FLOR (5)	6.2e+2 (3.6e+2 ... 1.2e+3)	2.4 (2.2 ... 2.5)	13 (11 ... 16)	1.6 (1.5 ... 1.6)
HAPPI-CCCMA happi2.0 (10)	-	-	15 (14 ... 17)	1.3 (1.2 ... 1.4)
HAPPI-ETH happi2.0 (10)	89 (48 ... 1.8e+2)	1.6 (1.4 ... 1.8)	11 (10 ... 13)	1.3 (1.2 ... 1.4)
HAPPI-NCC happi2.0 (10)	-	-	15 (13 ... 16)	1.6 (1.5 ... 1.7)
HAPPI-MIROC happi2.0 (10)	8.4 (5.4 ... 13)	0.87 (0.68 ... 1.1)	9.8 (8.6 ... 11)	1.3 (1.1 ... 1.4)
ACCESS ESM1-5 Historical+SSP245 (1)	1.1 (0.27 ... 4.7)	0.061 (-0.58 ... 0.68)	-	-
INM-CM4-8 Historical+SSP245 (1)	1.2e+2 (12 ... 1.9e+3)	1.4 (0.78 ... 2.0)	-	-
ACCESS-ESM1-5 (40)	2.1 (1.6 ... 2.6)	0.30 (0.20 ... 0.41)	3.0 (2.8 ... 3.2)	0.46 (0.43 ... 0.49)
BCC-CSM2-MR (1)	34 (4.4 ... 2.3e+2)	1.3 (0.54 ... 2.1)	10 (5.8 ... 18)	0.93 (0.70 ... 1.2)
CMCC-ESM2 (1)	12 (3.0 ... 55)	0.83 (0.38 ... 1.3)	17 (9.3 ... 33)	0.93 (0.76 ... 1.1)
EC-Earth3 (6)	8.3 (5.2 ... 13)	0.80 (0.62 ... 0.98)	4.7 (4.0 ... 5.5)	0.59 (0.53 ... 0.66)
EC-Earth3-CC (1)	4.1 (1.7 ... 9.7)	0.56 (0.21 ... 0.92)	4.1 (2.8 ... 6.2)	0.57 (0.41 ... 0.74)
EC-Earth3-Veg (7)	9.0 (6.0 ... 14)	0.82 (0.66 ... 0.98)	4.4 (3.8 ... 5.1)	0.57 (0.51 ... 0.63)
EC-Earth3-Veg-LR (3)	13 (5.6 ... 29)	0.95 (0.63 ... 1.3)	4.5 (3.5 ... 6.0)	0.61 (0.50 ... 0.72)
INM-CM4-8 (1)	7.0 (1.2 ... 36)	0.62 (0.058 ... 1.2)	8.6 (4.8 ... 15)	0.68 (0.49 ... 0.88)
INM-CM5-0 (1)	4.3e+2 (42 ... 4.3e+3)	1.6 (1.0 ... 2.1)	28 (14 ... 61)	0.99 (0.80 ... 1.2)
UKESM1-0-LL (5)	17 (9.3 ... 32)	0.96 (0.73 ... 1.2)	6.7 (5.5 ... 8.2)	0.68 (0.62 ... 0.74)
IPSL-CM6A-LR (32)	58 (36 ... 88)	1.4 (1.3 ... 1.6)	6.5 (5.5 ... 7.3)	0.94 (0.84 ... 1.0)
CNRM-CM6-1-HR HighResMIP (1)	2.3e+2 (9.9 ... 2.1e+4)	2.5 (1.1 ... 3.9)	-	-
HadGEM3-GC31-HM HighResMIP (1)	38 (2.1 ... 9.9e+2)	1.4 (0.31 ... 2.4)	-	-
HadGEM3-GC31-MM HighResMIP (1)	3.8e+2 (20 ... 2.6e+4)	2.4 (1.1 ... 3.7)	-	-

431 4.3 Hazard synthesis

432 Here we present the synthesis for PR and ΔI for the present compared to the 1.2°C cooler
433 climate (Fig. 5(a-b)) and the future vs. present (Fig. 5(c-d)). The best estimates of the PR range
434 from 1.1 in the ACCESS models to more than 600 in FLOR. The change in intensity goes from
435 almost no change to more than +2°C. Both are compatible with the highly uncertain
436 observational analysis; therefore, we can use the weighted mean to indicate the main result of
437 this study. Our synthesis concludes an event probability ratio of 30 (2 - 470) (Fig. 5(a)) and a
438 corresponding change in intensity of 1°C (0.2°C - 2.1°C) (Fig. 5(b)).

439 The change in PR for a further 0.8°C global temperature increase is 8 (3-12) (Fig. 5(c)) and an
440 additional increase in intensity of 1°C (0.3°C - 1.7°C) (Fig. 5(d)). The simulations based on
441 the HAPPI ensemble are centred at 1°C warming for the present day climate instead of 1.2°C
442 thus they show changes in likelihood and intensity for an additional 1°C of global warming
443 rather than 0.8°C. Nevertheless, the discrepancy between the individual models is smaller than
444 for the changes up until today.

445
446 Part of the uncertainties in the estimates from the participating models arise from differences
447 in the aerosol representation and the GHG and non-GHG forcings in the individual models.
448 For example, the PR and intensity change of the ACCESS-ESM1.5 model in Fig. 5 are
449 relatively low. This may be due to the relatively large global-mean aerosol indirect effect for
450 that model over the historical period that results in smaller historical warming as compared to
451 the other models in the CMIP6 ensemble considered in the study (Wang et al., 2021). The
452 differences in external forcing (SSP245 and SSP585) that are used to obtain the future scenario
453 runs in the participating models also contribute to uncertainties in PR and intensity changes for
454 the future (Fig. 5(c-d)). Furthermore, the choice of a Gaussian fit rather than a GPD might have
455 led to an underestimation of the changes in a relatively rare event such as this one
456 (Supplementary Section S3; Fig. S9). We therefore conclude that our overarching results are
457 conservative and the true influence of human-caused climate change is towards the higher end
458 of the estimated changes in likelihood. On the other hand, the central estimate of ΔI that the
459 2022 is made ~1°C warmer due to climate change is consistent across both observations and
460 models, and consistent with statements made in other dry regions (Daramola & Xu, 2022; Li,
461 Chen, & Li, 2019), thus lending high confidence to our findings.

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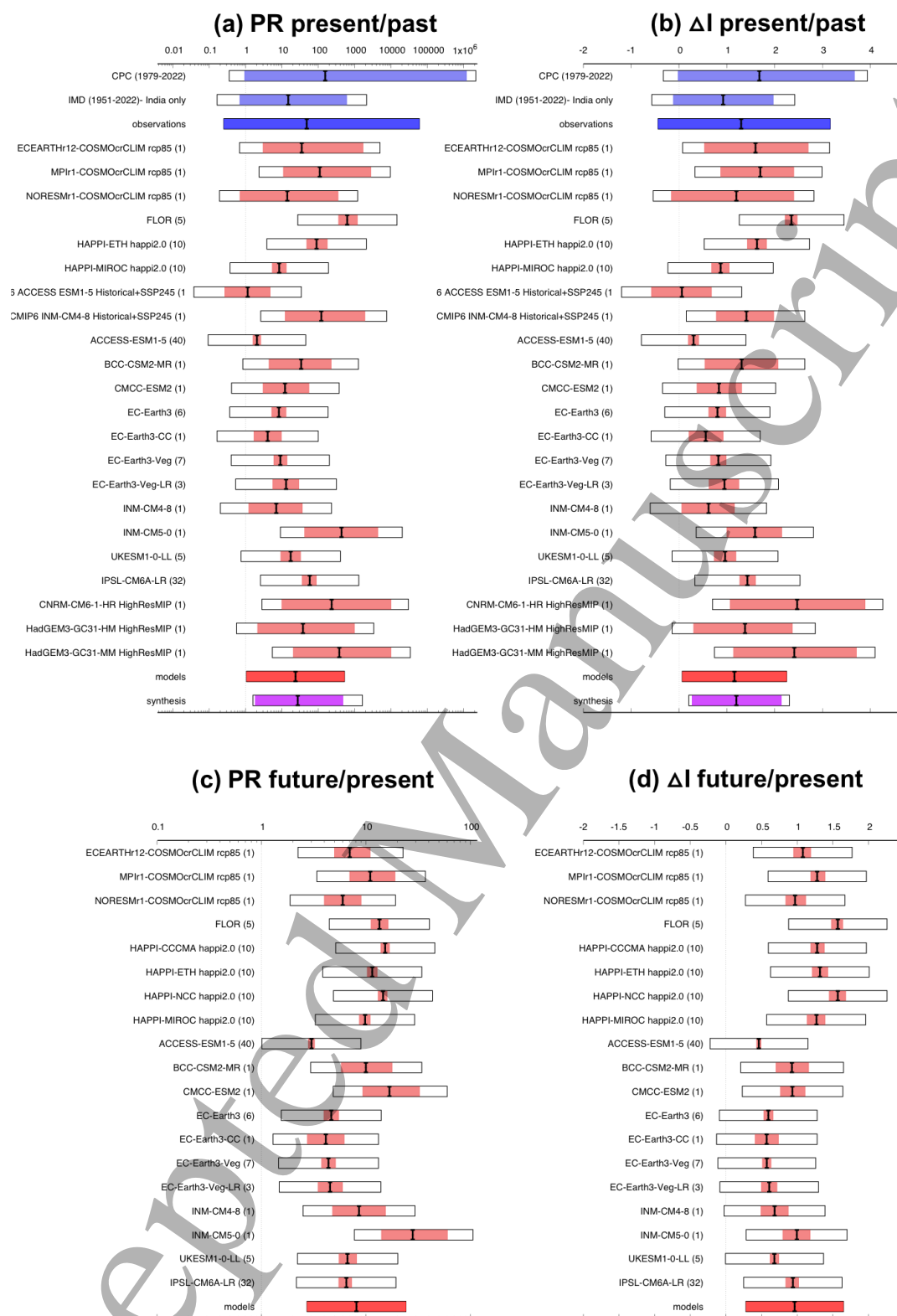


Fig. 5 (a) Synthesis of intensity changes (ΔI) and (b) probability ratios (PR) when comparing the 100-year event in today's climate with a 1.2°C cooler climate. (c) same as (a) and (d) same as (b) when comparing the 100-year event in today's climate with a 0.8°C warmer climate (equivalent to 2°C of global warming).

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5 477 **5 Vulnerability and exposure**

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8 478 The overall risk associated with the heatwave event is governed by the hazard as well as the
9 479 vulnerability and exposure (V&E) factors that make people and human systems more or less
10 480 susceptible to the impacts of the prolonged high temperatures. Heatwaves are “silent disasters”
11 481 because their impacts are difficult to ascertain. The 2022 heatwave is estimated to have led to
12 482 around 98 deaths in India and Pakistan (Datt, 2023; Irfan, 2023). However, heat-related deaths
13 483 are often undercounted across the globe; therefore the actual toll is likely higher (Ghumman &
14 484 Horney, 2016). A comprehensive review of the V&E factors that accompanied the 2022
15 485 heatwave is necessary for informing timely interventions and long-term strategies to address
16 486 vulnerability and improve preparedness for avoiding future impacts in these regions. We carry
17 487 out this review for five key areas revolving around demographics, informality in urban areas,
18 488 heat action planning and preparedness, agriculture and other compounding risks.

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21 489 5.1 Demographics and vulnerable groups

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24 490 The effects of the 2022 heatwave were primarily felt over northwestern India and Pakistan,
25 491 where some of the largest and densest urban areas in the world are situated (The Global
26 492 Statistics, 2022; ul Haque et al., 2021). Between 1983 and 2016, urban population exposure to
27 493 extreme heat is reported to have increased by approximately 200%, globally (Tuholske et al.,
28 494 2021). Three of ten cities that experienced the largest increase during this period are in our
29 495 study region, namely, New Delhi, Karachi, and Lahore.(Tuholske et al., 2021). Although
30 496 anybody can feel the impacts of extreme heat, vulnerable groups of people are affected
31 497 disproportionately. The most affected groups include outdoor workers such as farm workers,
32 498 labour migrants, low-income households, homeless people, daily wage earners, construction
33 499 workers, street vendors, street sweepers and rickshaw drivers (Climate & Development
34 500 Knowledge Network, 2016; Mazdiyasi et al., 2017), the elderly and young children, people
35 501 with chronic conditions (cardiovascular, respiratory, and cerebrovascular), people with pre-
36 502 existing mental illness, and people with cognitive and/or physical impairments (Carleton, 2017;
37 503 Mazdiyasi et al., 2017; Swain, Bhattacharya, Dutta, Pati, & Nanda, 2019). Tourists, travellers
38 504 and migrants are prone to additional risks due to missing warnings in local language, not
39 505 knowing how to access cool spaces, or being less accustomed to the local temperatures (Hari
40 506 et al., 2021).

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46 510 5.2 Informality in urban areas

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49 511 Approximately 10 million Karachi residents and half of New Delhi’s population live in
50 512 informal, low-income settlements (Pabani, 2021; World Population Review, 2022) with
51 513 building structures and roof types that significantly intensify indoor temperatures during the
52 514 day (Mahadevia et al., 2020; Mukhopadhyay et al., 2021). In the absence of adequate cool roof

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3 515 retrofits(Vellingiri et al., 2020), the urban poor rarely get respite from the extreme heat (Weitz
4 516 et al., 2022), especially those who spend most of their time indoors, such as the elderly, women,
5 517 and people with physical impairments. The elderly low-income residents are up to 4.3 times
6 518 more likely to be exposed to hazardous heat than their rural counterparts (Weitz et al., 2022).
7 519 Moreover, tin roofs in such settlements are known to further exacerbate the urban heat island
8 520 effect (Padmanaban, 2021). Half of India's workforce is estimated to be outdoor labourers
9 521 (Jha & Kishore, 2022). Therefore, it is not surprising that India faces the largest impacts of heat
10 522 on heavy manual labour such as agriculture and construction, with over 101 billion working
11 523 hours lost per year (out of the global sum of 228 billion; Parsons et al., 2021). Under future
12 524 warming, both India and Pakistan are amongst the top ten countries projected to experience the
13 525 largest population-weighted labour losses, together with China, Bangladesh, Indonesia, Sudan,
14 526 Vietnam, Nigeria, Thailand and Philippines (Parsons et al., 2021).

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20 527 The urban poor — in particular, daily wage earners who work outside are the worst-off during
21 528 heatwaves, due to the effects of direct sunlight exposure being compounded by air pollution,
22 529 limited access to healthcare facilities and inadequate mitigation practices (Anwar et al., 2022;
23 530 Bakhsh et al., 2016; Barthwal et al., 2022). At least four people found dead on the streets in the
24 531 city of Nagpur in India, in April 2022, were suspected to be heat stroke-related fatalities
25 532 (Mascarenhas, 2022). The urban heat island effect, which can exacerbate heat by up to 12°C
26 533 locally, also contributes to the exposure risk among the urban poor (Razzak et al., 2022). For
27 534 peri-urban residents, the risk of exposure is largely due to the long-distance commute to work
28 535 by foot, two-wheelers or public transportation, thus limiting their options to mitigate (Bakhsh
29 536 et al., 2016). Finally, it is important to note that this group had already been reeling under the
30 537 effects of the COVID-19 pandemic (Raju et al., 2021) before the 2022 extreme heat episode.
31 538 The impacts of the heatwaves such as the 2022 event could therefore make the pandemic
32 539 recovery even longer, highlighting the need for anticipatory humanitarian approaches
33 540 (Thalheimer et al., 2022).

34 541 5.3 Heat Action Planning, Preparedness, and Response

35 542 Given the propensity for heatwaves in the populous South Asian regions, the countries have
36 543 implemented an arsenal of early warning systems and early action programmes at local to
37 544 regional scales for mitigating impacts (Das & Smith, 2012; Vahlberg et al., 2022). For example,
38 545 the South Asia Heat Health Information Network (SAHHIN;
39 546 <https://climateandcities.org/about-us/south-asia-heat-health-information-network/#>) was
40 547 developed in 2020 to share lessons and increase capacity to deal with extreme heat across South
41 548 Asia. Both India and Pakistan are making significant and rapid strides to combat extreme heat
42 549 in particular, especially in recent decades.

43 550 5.3.1 Heat Action Planning, preparedness, and response in India

44 551 In the aftermath of the catastrophic heatwave in 2010, Ahmedabad in India became the first
45 552 South Asian city to implement a Heat Action Plan (HAP); the city is now estimated to avoid
46 553 approximately 1,190 heat-related deaths annually (Hess et al., 2018). Since then, over 120
47 554 Indian cities and states have developed HAPs that focus on building public awareness and

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3 555 capacity among health professionals, issuing safety alerts for residents, fostering inter-agency
4 556 coordination, and enabling adaptive measures for vulnerable groups that include adopting cool
5 557 roof technology, increasing green coverage in urban areas, assembling roofing structures at
6 558 markets, and installing drinking water stations along highways (Natural Resources Defense
7 559 Council, 2022; Padmanaban, 2021). Aimed at mitigating the increasing temperatures' toll on
8 560 public health and building consensus around its management, India's Ministry of Health and
9 561 Public Welfare (with support from other government departments and non-governmental
10 562 actors) developed the National Action Plan on Heat Related Illnesses (National Centre for
11 563 Disease Control, 2021). Launched in 2021, it contains guidelines for the government, health
12 564 care facilities and policymakers on managing and reporting heat-related illnesses. Stocktaking
13 565 basic equipment and medicine and ensuring sufficient staffing are some of the recommended
14 566 actions when faced with extreme heat.

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20 567 In anticipation of the 2022 heat season, the National Disaster Management Authority (NDMA)
21 568 held a national workshop on heat preparedness, mitigation and management, in March (Natural
22 569 Resources Defense Council, 2022). In a first, the IMD implemented an impact-based early
23 570 warning system, providing accessible and actionable information for increasing heat risk
24 571 preparedness and coping capacity (Natural Resources Defense Council, 2022). The agency has
25 572 been disseminating timely information to the public since 2020 via their mobile phone
26 573 application "Mausam". Thus, there was an increased awareness of the weather and warnings
27 574 this year (Natural Resources Defense Council, 2022). Bulletins from the Ministry of Health
28 575 and the Indian Institute of Public Health Gandhinagar (IIPHG) advised people to wear
29 576 lightweight clothing of natural fibres, avoid exposing one's head to direct sunlight and seek
30 577 care if they recognize any signs of heat-related illness (PTI, 2022). Preparing for a heavy
31 578 inflow, hospitals across India set up special wards for heat-related illnesses, rolled out capacity-
32 579 building training and sensitisation on heat risk and symptoms for medical staff, and were
33 580 instructed to ensure uninterrupted electricity supply to guarantee the functioning of cooling
34 581 devices (Mascarenhas, 2022; N. Singh, 2022; TN National Desk, 2022). Cooling centres and
35 582 rooms were established in primary health centres, hospitals, places for worship, malls and other
36 583 public buildings to provide visitors with drinking water, health care and respite from the heat,
37 584 while fans and cooling structures were installed in schools (Lal, 2022).

38 585 *5.3.2 Heat Action Planning, preparedness, and response in Pakistan*

39 586 The Start Network - a conglomeration of agencies from across the world for aiding
40 587 humanitarian action, has a national disaster risk financing programme for Pakistan
41 588 (<https://startnetwork.org/disaster-risk-financing-pakistan>) that funds early action in
42 589 anticipation of heatwaves. Activities include training community leaders in disaster
43 590 preparedness and first aid, opening shelters in schools and other communal spaces, spreading
44 591 public awareness on heatstroke prevention and identification of symptoms, establishing
45 592 helplines, and setting up health emergency camps that provide cold drinking water and
46 593 medicines (Start Network, 2021). Cities covered by the programme include Karachi, Larkana,
47 594 Multan, Sibi, Nawabshah and the city of Jacobabad which incidentally recorded the region's
48 595 maximum temperature during the 2022 heatwave, at 49°C on 30 April 2022 (Bhatti, 2022).

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3 596 Following the devastating 2015 heatwave that led to over 1,200 deaths Pakistan (Rafferty,
4 597 2016) with more than 65,000 heatstroke-related hospitalisations in Karachi alone (Glum,
5 598 2015), Karachi and other urban areas across Pakistan have developed HAPs (Commissioner
6 599 Karachi, n.d.) that outline the immediate actions following a heatwave warning, such as
7 600 establishing cooling centres in places for worship, malls and other public buildings, increasing
8 601 staffing at healthcare centres to accommodate a rise in patient influx and redistributing more
9 602 ambulances to densely populated areas (Commissioner Karachi, n.d.).
10 603 Adaptive measures such as increased water consumption, staying in the shade or bathing more
11 604 frequently, is paramount for reducing heat-related mortality in urban Pakistan (Bakhsh et al.,
12 605 2018). A Start Network evaluation on early action in response to a 2021 heatwave in the city
13 606 of Sibi showed that most people tend to apply these strategies, except those for whom it would
14 607 negatively impact livelihoods, such as rickshaw drivers or construction workers (Guyatt &
15 608 Khan, 2022). This dilemma to choose between safeguarding one's health and sustaining one's
16 609 livelihood is characteristic of the most at-risk populations' exceptional vulnerability. In
17 610 October 2021, as Pakistan updated its Nationally Determined Contributions (NDC), the
18 611 government announced that it is developing a Cooling Action Plan to be adopted by 2026
19 612 (United Nations Climate Change, n.d.). The plan will identify key cooling needs and outline
20 613 sustainable actions for addressing those needs, both current and prospective.

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28 614 In response to the 2022 heatwave, public health authorities in Pakistan instructed health units
29 615 to open "heatstroke centres" to help the public connect with the authorities, while also
30 616 reminding people to avoid direct sunlight and increase their water consumption (A. Saeed,
31 617 2022; Toheed, 2022). Although the most rigorous action seems to have been taken in May
32 618 (Web Desk- Geo News, 2022), numerous trainings were rolled out in April. Between 18 and
33 619 29 April, the Provincial Disaster Management Authority (PDMA) Sindh and Pakistan Red
34 620 Crescent Society (PRCS) jointly offered heat emergency training to traffic police and line
35 621 department officials as well as representatives of civil society organisations.

36 622 5.4 Agriculture

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42 623 The agriculture and related sectors are the major contributors to the national economies of India
43 624 and Pakistan, with 60% and 40% of the respective population working in this sector (Statista,
44 625 2022b, 2022a). The 2022 heatwave hit at a critical time, during the final period of the growing
45 626 season for winter crops such as wheat and barley, and also affected summer crops such as
46 627 pulses, coarse cereals, oilseeds, vegetables and fruits.

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50 629 India, the second-largest wheat producer globally, is also a major consumer. Farms in the
51 630 northern states of Punjab and Haryana of India that account for 25% of the country's total
52 631 wheat production (United States Department of Agriculture, 2022) and Uttar Pradesh lost an
53 632 estimated 10%-35% of crop yields due to the heatwave (Ghosal, 2022), affecting local market
54 633 prices, that rose to 15% in some regions (Arora & Bhardwaj, 2022). Global food prices also
55 634 reached their highest level ever recorded in March 2022, with a 40% rise since the beginning
56 635 of the year (FAO Food Price Index, 2022), due to the Russian invasion of Ukraine, a major
57 636 wheat producer (Parija & Bhatia, 2022) and increasingly high fertiliser prices this year (Meyer,
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2022). Therefore, the Indian government was forced to retract its initial plan to boost its wheat exports to meet the global wheat shortages and to impose a ban on export, to protect India's internal food market — further affecting the global wheat market and food-dependent countries (Hoskins, 2022). In Pakistan, exportable mango varieties have seen a 50% loss and 30% in local varieties, due to the extreme heat, which was followed by a pest attack, with yields per acre falling from 40 to 28 maunds (1 maund=37.32kg) (Ilyas, 2022).

Although advisories were sent to farmers to ensure frequent irrigation for the crops (e.g., [Government of India, 2022](#)), functional electricity and water systems are also important during periods of extreme heat. At present, there is an urgent need for research, public policies and investments to focus on adaptation strategies to minimise the future impacts of extreme heat on agriculture.

5.5 Compounding risks

In addition to the direct impacts of extreme heat on public health, agriculture, socio-economic factors and urban planning discussed above, there are compounding risks such as cascading hazards and energy availability. Heatwaves are known to create cascading hazards, leading to secondary events of significant impact (Pescaroli & Alexander, 2014; Tilloy et al., 2019; Vogel et al., 2020). For example, increased temperatures and evapotranspiration from heatwaves can result in both water shortages and floods from meltwater. Spikes in energy demand during heatwaves can result in shortages, thereby limiting means for cooling and irrigation.

In northern Pakistan and India, rapidly melting glaciers are putting thousands at risk of glacial lake outburst floods (GLOFs) and landslides as well as to decreased water supplies. GLOF risks were highlighted by the Pakistani government in their heatwave response (Government of Pakistan, 2022) and a large one occurred on 7 May 2022, wiping out a bridge, houses and inundating farmland in the Hunza valley (Davies, 2022). Heatwaves also increase the risk of forest-fires (Jain et al., 2021). On April 27th, the Forest Survey of India reported 300 active large forest fires, a third of which were in the Uttarakhand province (NASA Earth Observatory, 2022). In Delhi, a massive landfill caught fire for at least 9 days (Express News Service, 2022). Across Pakistan, multiple farm and village fires have been reported throughout April, resulting in loss of lives and properties (Provincial Disaster Management Authority, 2022; The Third Pole, 2022). In turn, these fires have a significant impact on air quality, which increases morbidity and mortality of extreme heat events. April was reported as the worst month for air quality in Delhi since 2015 - the city recorded 29 days of "poor air quality" (200-300 Air Quality Index, AQI) (Paljor, 2022). Throughout March and April, Lahore consistently measured AQI corresponding to levels "unhealthy for sensitive groups" (151-200) and "unhealthy" (201-300) (Environment Protection Department, 2022). About 70% of India's electricity generation comes from coal (IEA, 2021), with about 60% of energy provision from coal, oil and natural gas in Pakistan (IEA, 2020). The 2022 heatwave increased the demand for coal imports in India due to shortages resulting in rolling blackouts (Chaturvedi, 2022). At least 16 out of 28 states in India experienced power outages of two and ten hours duration (Bloomberg, 2022), affecting the public, industry, and agriculture (R. K. Singh, 2022).

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6. Concluding remarks

682 In 2022 India and Pakistan experienced an intense heatwave that began in early March and
683 persisted into the month of May. Given the devastating impacts on human health and
684 agriculture we set out to answer the question of whether and to what extent the event could
685 have been influenced by climate change. Therefore, we performed an attribution analysis of
686 the observed March–April average daily maximum temperature, using published peer-
687 reviewed methods. Upon comparing the event characteristics — return period and intensity, in
688 today’s climate with counterfactual worlds without climate change (GMST 1.2°C cooler as
689 compared to now), the 2022 heatwave is found to be made 1°C hotter and 30 times more likely
690 by climate change. Notwithstanding the conflicting effects of local factors on magnitudes such
691 as aerosol interactions among the different models, future projections also show consistently
692 more intense heat waves of longer durations and occurring at a higher frequency over India
693 (Murari et al., 2015; Mishra et al., 2017) and Pakistan (Saeed et al. 2017; Nasim et al., 2018).
694 Our results show that under global warming of 2°C above pre-industrial levels, the 2022
695 heatwave is expected to become 2–20 times more likely and 0.5°C–1.5°C hotter than now.

696 The urban poor in India and Pakistan are amongst the most exposed and vulnerable to extreme
697 heat, and are left using coping mechanisms to withstand the extreme heat and earn a daily wage.
698 Rising temperatures from more intense and frequent heatwaves will render coping mechanisms
699 inadequate, as some regions meet and exceed limits to human survivability (Mora et al., 2017).
700 While some losses will inevitably occur due to extreme heat, it is misleading to assume that
701 the impacts are inevitable (Raju et al., 2022). This emphasises the need to record losses and
702 damages occurring due to climate change related disasters (Boyd et al., 2021). Adaptation to
703 extreme heat has been shown to be effective in some cases (Hess et al., 2018). Heat Action
704 Plans that include early warning and early action, awareness raising and behaviour changing
705 messaging, and supportive public services can reduce mortality, and India’s rollout of these
706 has been remarkable, now covering 130 cities and towns. There are, however, still large
707 research gaps on adaptation to heat across India and Pakistan that will require further study to
708 build a stronger evidence base for action (Pachure et al., 2022). Heatwaves such as the event
709 we analysed here are considered disasters due to people’s vulnerabilities, an issue that needs
710 to be tackled by society. Better urban and health planning, disaster insurances and livelihood
711 protection mechanisms, investment in green spaces, energy grid strengthening, improved water
712 infrastructure and pollution controls could all contribute to ensure that fewer people suffer as
713 temperatures rise.

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Data Availability Statement

716 The data that support the findings of this study are openly available at
717 <https://climexp.knmi.nl/HeatwaveIndiaPakistan2022.cgi>

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