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Origin and evolution of the genera *Lonsdaleia* and *Actinocyathus*: Insights for the Mississippian palaeogeography from the western Palaeotethys



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Abstract Representatives of the subfamily Lonsdaleiinae Chapman, 1893 are common in the Mississippian of the western Palaeotethys. A general analysis of the origin, distribution and evolution of that subfamily has been undertaken. The most probable hypothesis for the origin of the genus *Lonsdaleia* McCoy, 1849 is to acquire colonialism via the genus *Axophyllum* Milne Edwards and Haime, 1851. *Actinocyathus* d'Orbigny, 1849 would be a descendant of *Lonsdaleia* by increasing integration in the colonies. The first occurrences of *Lonsdaleia* have been recorded in the lower Visean from northern Britain and northern Tianshan Mountains of northwestern China, but the diversification and migration to the whole Palaeotethys only happened in the late Visean. Three hypotheses are proposed on that matter. The Serpukhovian was also a period of migrations and diversification for these genera. Both *Lonsdaleia* and *Actinocyathus* have been recorded in Bashkirian refuges, the Sverdrup Basin in northern Laurasia and the Tindouf Basin in northern Africa, respectively. The division of the western Palaeotethys into six subprovinces based on the distribution of corals is proposed.

Keywords Carboniferous, Palaeotethys, Biogeography, Evolution, Rugosa

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The subfamily Lonsdaleiinae Chapman, 1893 comprises colonial corals of the family Axophyllidae (Hill, 1981). The genera *Lonsdaleia*, *Actinocyathus* and *Serraphyllum* are included in that subfamily. The first two genera are abundant and almost cosmopolitan. The third one has been recorded only in the Montagne Noire up to now (Poty and Hecker, 2003). Some other genera have been also included in the subfamily, but they have been regarded as synonyms of the cited ones (Hill, 1981).

The Lonsdaleiinae have been cited in most areas of the world in the upper Visean and the Serpukhovian (see references below), and more rarely in the Bashkirian. However, their distribution is guite irregular for several reasons: The knowledge of the Mississippian coral record shows strong variations in different geographical areas. In addition, the name Lonsdaleia has been used for Pennsylvanian corals that belong to other families such as Petalaxidae or Waagenophyllidae (Dobrolyubova, 1936; Douglas, 1936; Easton, 1960; De Groot, 1963). Moreover, some citations may not be consistent, either because the identification is not clear, or because the age of the occurrence is doubtful (Grosch, 1912; Hudson, 1958). An extensive revision of the occurrences of the genera Lonsdaleia and Actinocyathus in different regions of the world provides new data that are useful to present a general view of the origin, evolution and migrations of these two genera.

The taxonomic position of the two involved genera, Lonsdaleia McCoy, 1849 and Actinocyathus d'Orbigny (1849), has been matter of discussion for a long time. Martin (1809) described and figured corals with the names Erismatolithus Madreporites (duplicatus) and Erismatolithus Madreporites (floriformis). Both were later described and figured under different names (see Smith, 1915 for a complete review of the older literature). McCoy (1849) defined the genus Lonsdaleia for fasciculate corals having a complex axial structure and lonsdaleoid dissepiments. d'Orbigny (1849) introduced the genus Actinocyathus for corals having massive (cerioid) habit, complex axial structure and lonsdaleoid dissepiments. McCoy (1851) established the species L. duplicata as the genotype of Lonsdaleia and used the name Strombodes for the massive corals with identical inner features. Milne-Edwards and Haime (1850-54) included both massive and fasciculate forms in a single genus that they called Lithostrotion (in 1850) and later Lonsdaleia (in 1851 and 1854). Fromentel (1861) introduced the name Stylidophyllum for the massive forms. This proposal was not much followed, because most authors in the subsequent

years used the name *Lonsdaleia* for both massive and fasciculate corals (Thomson and Nicholson, 1876; Vaughan, 1905; Garwood, 1912; etc.), but some authors used that name (Chi, 1931; Gorsky, 1935).

Lang et al. (1940) stated that the name Actinocyathus is available. Kato (1966) demonstrated that it has priority over Stylidophyllum Fromentel (1861) that was used for cerioid Lonsdaleiinae by some authors (Gorsky, 1935; Dobrolyubova, 1936, 1958; Yamagiwa, 1961), either as a subgenus of Lonsdaleia or as a generic name. Kato (1966) proposed that Actinocyathus should substitute Stylidophyllum if the latter could be considered as a separate genus. Thus, Sando (1975) postulated it as a subgenus of Lonsdaleia. Hill (1981) proposed the use of Actinocyathus as a separate genus from Lonsdaleia. Subsequently, some authors considered it as a subgenus (Poty, 1981; Sando, 1983; Poty and Hecker, 2003; Hecker, 2012) and others as a genus (Mitchell, 1989; Wang, 1989; Rodríguez et al., 2013a). The differentiation of genera based on different habits is a common feature in Paleozoic corals (e.g. Lithostrotion Fleming, 1828, Siphonodendron McCoy, 1849), but there are different opinions between the specialists on rugose corals on that matter (Hill, 1938-41, 1981; Fedorowski, 1978, 1981, 1984; Poty, 1981, 2010). In the case of Lonsdaleia and Actinocyathus, as there is no consensus, and since it is a matter of subjectivity, we will use both names as separate genera, because it simplifies the terminology.

2. Discussion

The Lonsdaleiinae from the Tindouf Basin in Morocco provide important information for the understanding of the Visean and Serpukhovian palaeogeography of the western Palaeotethys, because they complete the view of the evolution of these corals in the context of Europe and North Africa. In order to analyse the distribution of the Lonsdaleiinae, it is necessary first to revise the origin and evolution of the genus Lonsdaleia s.l. (including the genera/subgenera Lonsdaleia and Actinocyathus).

2.1. Lonsdaleia origin and diversification

Several hypotheses have been proposed in order to explain the origin of the genus *Lonsdaleia*. Thomson (1883) suggested that *Thysanophyllum* (a massive genus similar to *Lonsdaleia*, but without or with reduced axial structure) could be the ancestor of *Lonsdaleia*. Carruthers in Garwood (1912) followed that hypothesis by comparing the neanic stages of *Lonsdaleia* and *Thysanophyllum*. Smith (1915), in his monograph on *Lonsdaleia*, supported it, showing the early neanic stages of these genera in his plate XVII. Vaughan (1905) proposed a different hypothesis when he suggested that *Lonsdaleia* was derived from the solitary genus *Clisiophyllum* by development of colonialism and lonsdaleoid dissepiments. Only a year later, Matley and Vaughan (1906) changed his hypothesis and proposed the genus *Carcinophyllum* as the ancestor of *Lonsdaleia*. The advantage of this genus for the ancestry is its lonsdaleoid dissepimentarium, and thus requiring only the development of colonialism for the appearance of the new genus.

Hill (1938-41) suggested that the relationship between Lonsdaleia and Thysanophyllum could be inverse, with the first being the ancestor of the latter. That hypothesis was followed by Vassiljuk (1960) and other authors, but they did not explain the ancestry of Lonsdaleia. Poty (1981) suggested that Axophyllum should be the ancestor of Lonsdaleia, following the hypothesis of Matley and Vaughan (1906), as Semenoff-Tian-Chansky (1974) demonstrated that Carcinophyllum is a synonym of Axophyllum. That hypothesis has been followed later by most authors (Poty and Hecker, 2003; Poty, 2010; Rodríguez and Somerville, 2010; Somerville and Rodríguez, 2010). Smith (1915) also suggested that the massive forms of Lonsdaleia derived from the fasciculate ones. He used the name Lonsdaleia, but now the name Actinocyathus is accepted for massive forms, either as a separate genus, or as a subgenus. Also, the consideration of Lonsdaleia as an ancestor of Actinocyathus is agreed by most authors (Hecker, 1997; Poty and Hecker, 2003; etc.).

An important question for the analysis of the distribution of the genera Lonsdaleia and Actinocyathus is when and where they appeared. The oldest Lonsdaleia species is L. praenuntia Smith (1915). The type specimens were collected by Vaughan (1911) in the lower Visean from Arnside, NW England. This species shows a reduced axial structure that resembles that of Cystolonsdaleia Fomichev, 1953 or Dorlodotia Salée, 1920. In fact, this species was not included in the biostratigraphical chart of Mitchell (1989), who considered the first appearance of Lonsdaleia to be much higher, in his zone H (Brigantian). The doubts on the identification of the specimens of Vaughan (1911) as Lonsdaleia are due to the absence of buddings that demonstrate its coloniality. Thus, they could also be regarded as representatives of Axophyllum. Poty (1980, 1981) suggested that this species could be the ancestor of the genus Dorlodotia. Hecker (2011, p. 47) indicated that "studies of variability in Dorlodotia suggest that Lonsdaleia praenuntia is also a Dorlodotia". We can't agree with this statement, because even if the axial structure of *L. praenuntia* is "loosely constructed and asymmetrical ..." as pointed out by Smith (1915, p. 243), the structure of the tabularium is typical of Lonsdaleia and, as noted by Smith (1915, p. 243), "the central column exhibits no marked difference from the same structure in L. duplicata." Fig. 1 shows the similitudes and differences: a complex axial structure and concave periaxial tabulae in *L. praenuntia* and simple axial structure and periaxial tabulae elevated towards it in Dorlodotia briarti. Consequently, Lonsdaleia praenuntia must be regarded as an early representative of that genus.

It is a common fact in many rugose corals that colonial taxa descend from solitary corals (Fedorowski, 1978; Poty, 2010; Rodríguez and Somerville, 2010; Denayer and Webb, 2015). In many cases, there are some intermediate stages of gregarism and of solitary corals developing offsets in the calice without further growth of a true colony. It has been called quasicolonialism (Fedorowski, 1978) or protocolonialism (Somerville and Rodríguez, 2010). As stated above, *Lonsdaleia* probably derived from a solitary species of *Axophyllum*, a genus which shares most inner features with it. A case of a gregarious and quasi-colonial axophyllid (*Howthia* Somerville and Rodríguez, 2010) was



Fig. 1 Comparison between *Dorlodotia briarti* (A, B, transverse and longitudinal sections; figures taken from Poty, 1981) and *Lons-daleia praenuntia* (C, D, transverse and longitudinal sections; figures taken from Smith, 1915). The main differences are (i) that the axial structure is more complex in *L. praenuntia*; it is joined to the counter septum in *D. briarti* and to the cardinal septum in *L. praenuntia*, and (ii) the periaxial tabulae are concave and more or less horizontal in *L. praenuntia*, but in *D. briarti* there are no axial and periaxial series of tabulae and they are elevated towards the medial plate.

described in the upper Tournaisian from Howth (Ireland). It could be an example of the way for the appearance of *Lonsdaleia* from *Axophyllum*. And it could be the ancestor of all the Lonsdaleiinae if they are not polyphyletic.

The main difficulty to accept *L. praenuntia* as the ancestor of the late Visean species of Lonsdaleia is the total absence of the genus in the middle Visean. However, several species have been recorded in the upper Visean of the same domain (the Alston block of northern Britain) (Smith, 1915). This could be due to two different reasons (1) L. praenuntia is not the ancestor of the late Visean species, and they evolved quickly during the Asbian transgression. The development of that species could be an early attempt to develop colonialism in the family, without descendants, or it could be an Axophyllum. (2) The presence of several species in the same area indicates a previous diversification during the middle Visean that is not recorded, or they evolved quickly during the late Visean.

However, the development of colonialism could have occurred more than one time in the Axophyllidae and Lonsdaleia could be polyphyletic. In many species of Lonsdaleia, the axial structure is closer to that of the genus Dibunophyllum than to that of Axophyllum. In Dibunophyllum and many species of Lonsdaleia, the axial structure is composed of a median lamella, a small number of radial lamellae and tabulae inclined to the periphery. The lamellae are usually anastomosed or winding. On the contrary, in Axophyllum, the axial structure is composed of thickened, more irregular and anastomosing radial lamellae, crossed by a median plate. This was the reason why Smith (1915) also suggested the possibility of that genus as the ancestor of Lonsdaleia. He rejected that possibility, but the similarities are conspicuous, because in many species of the Lonsdaleia duplicata and Actinocyathus floriformis groups, the axial structure is mainly dibunophylloid. We checked many specimens of Lonsdaleia from different provenances (Tindouf, Betic Cordillera, Britain, Carnic Alps, Moscow Basin, etc.) and all of them show a consistent homogeneity in the microstructure that fit with the Axophyllidae family (lamellar wall, tabulae, dissepiments and septal stereoplasm), but not with Dibunophyllum (granulofibrous mesoplasm and fibrous stereoplasm).

Actinocyathus and Lonsdaleia can be divided into several groups of species, with different features, and probably with different ecological constraints that conditioned their palaeogeographic distribution. Hecker (1997, 2010) gave a complete view of the species groups of Actinocyathus; the species group of A. floriformis Martin with short minor septa and the species group of *A. crassiconus* with long minor septa (Fig. 2). As she explained in detail their geographic and ecological distribution, we will not develop this matter here. Some species belonging to the *A. floriformis* group are *A. borealis*, *A. bronni* and *A. rossica* (Hecker, 1997). Typical representatives of the *A. crassiconus* group are *A. gorskyi*, *A. latevesiculosus*, *A. mariae*, *A. sarytschevae*, and *A. subtilis*.

In Lonsdaleia, the type species Lonsdaleia duplicata and most species show a dibunophylloid axial structure and thin structures and lived mainly in stable carbonate platforms. Some of the species included in this group are L. alstonensis Smith (1915), L. arctica Gorsky (1935), L. caledonia Smith (1915), L. crassigemmata Dobrolyubova (1958), L. duplicata Martin (1809), L. grandicaspia Degtiarev (1965), L. majiobaensis Fan (1978), L. melmerbiensis Smith (1915), L. multiseptata Dobrolyubova (1958), L. permanoseptata Vassiljuk (1960), L. siblyi Smith (1915), L. singularis Dobrolyubova (1958), and others. In contrast, some other species that we will name the Lonsdaleia corbarensis group, show smaller corallites, thick structures and more irregular axophylloid axial structure. The species that we include in this group are L. agapoviensis Kachanov, 1964, L. corbariensis Semenoff-Tian-Chansky and Ovtracht (1965), Lonsdaleia carnica Rodríguez et al., 2019, L. redondensis Poty and Hecker



Fig. 2 Inferred phylogenetic relationships between the different species groups in Lonsdaleiinae. Abbreviations: *A. = Actinocyathus; L. = Lonsdaleia;* gr. = Group.

(2003), L. reutheri Boll (1985), L. taveli Altmark (1963) and L. tichyi Dobrolyubova (1958). Most of these species lived in more unstable environments, associated to microbial communities, and their geographical distribution differs greatly from those of the Lonsdaleia duplicata group. Poty and Hecker (2003) demonstrated the presence of a biform tabularium and periaxial cones in several species of Lonsdaleia. We checked all our specimens and many in the literature to analyse the presence of these features in the mentioned groups. In many cases we couldn't find proofs of their presence because the published figures don't show these features, but they seem to be present in several species of both groups of Lonsdaleia and in Actinocyathus (Table 1).

2.2. Lonsdaleia and Actinocyathus distribution

Any palaeogeographic analysis needs to take into account the plate tectonics, and the areas to be

compared should be consistent with the different terranes that are known in each period of geological history. We tried to follow this axiom in the analysis of the distribution of the genera Lonsdaleia and Actinocyathus, but sometimes, the record itself or the studies in some areas are scarce. For example, in Germany, three palaeogeographic domains are recognised for the Carboniferous: the Rhenohercynian, the Saxothuringian and the Moldanubian, from north to south (Wever, 2000), but records in those domains are very scarce and in most cases they occur in allochthonous blocks or in reworked facies (olistostromes or culm). Thus, for this study, these three domains will be amalgamated as one single area (Germany). The same has been done with other areas, such as the three main Austrian outcrops yielding Mississippian corals (Carnic Alps, Nötsch and the Greywacke Zone; Hubmann, 2002) that will be considered as one unit. We also use the name Poland for several palaeogeographic units in that country (Sudetes, Upper Silesian Basin, Lublin Basin, and also the prolongation southeastwards

Table 1 Comparison of biometric data for species of *Lonsdaleia* and *Actinocyathus* mentioned in this paper. Abbreviations: N = Number of major septa; Dt = Diameter of tabularium; Das/Dt = Ratio of diameter of axial structure to tabularium diameter; RL = Number of radial lamellae; B.T. = Biform tabularium; ? = Not determined.

Species	Ν	Dt	Das/Dt	RL	B.T.	Minor septa
Lonsdaleia						
L. duplicata Martin	25–27	4.5–6	1/3-1/2	10–20	Present	Short
L. agapoviensis Kachanov	25–27	9.0-11.0	1/4	5—10	No	None
L. alstonensis Smith	26–28	4.0-5.0	1/2	15—30	?	Short
L. arctica Gorsky	22–24	3.0-5.0	1/3-1/2	20-30	?	Short
L. caledonia Smith	28-30	7.0-9.0	1/2	15—25	No	Long
L. corbariensis Semenoff-Tian-Chansky	20—25	5.0-7.0	1/2	15–20	?	Long
L. crassigemmata Vassiljuk	25-30	7.0-9.0	1/5-1/4	6—10	No	Short
L. grandicaspia Degtiarev	30-35	10.0-12.0	1/3-1/2	25-30	?	Long
L. majiobaensis Fan	36-37	10.0-12.0	1/4	20-25	Present	Short
L. melmerbiensis Smith	26–28	6.0-8.0	1/3	5-10	Present	Short
L. multiseptata Dobrolyubova	27–28	9.0-10	1/3	15—25	Present	Long
L. permanoseptata Vassiljuk	18—21	5.0-7.0	1/3	13—15	?	Short
L. praenuntia Smith	26-28	5,0-6,0	1/3	5—10	?	None
L. redondensis Poty and Hecker	15—18	3-4.2	1/3	5—10	Present	Long
L. reutheri Boll	24	7.0-10.0	1/3	5—10	Present	Short
L. <i>siblyi</i> Smith	22–24	5.5-6.5	1/3-1/2	10–15	No	Short
L. singularis Dobrolyubova	27–29	9.0-10.0	1/3-1/2	15—25	?	Short
L. taveli Altmark	27–28	6.0-9.0	1/3	8—10	?	Long
L. tichyi Dobrolyubova	20-23	5.0-7.0	1/3	5-10	?	None
L. carnica Rodríguez	22–25	4.3-5.2	1/4-1/3	5—20	No	Long
Actinocyathus						
A. borealis (Dobrolyubova)	19–23	5.0-6.5	1/4-1/2	10—20	Present	Short
A. bronni (Milne-Edwards and Haime)	30-34	7.0-10.0	1/3	10—15	No	Short
A. crassiconus (McCoy)	22–27	4.8-6.5	1/2	15—30	Present	Long
A. floriformis (Martin)	20—25	4.0-6.0	1/3	5-10	Present	Short
A. gorskyi (Dobrolyubova)	26-33	5-7.5	1/2	15–24	Present	Long
A. latevesiculosus (Dobrolyubova)	22–27	6-8.5	1/3	7–12	Present	Long
A. mariae Rodríguez et al.	23–28	6—9	1/3-1/2	10-30	Present	Long
A. rossicus (Stuckenberg)	19—24	5.0-6.0	1/3-1/2	15—25	No	Short
A. sarytschevae (Dobrolyubova)	34-40	7.5-10.3	2/3	10-30	Present	Long
A. sp. A Hecker	21-26	4.8-6.5	1/3-1/2	8–20	Present	Long
A. subtilis (Dobrolyubova)	21-30	4-6.3	1/3-1/2	10-24	Present	Long



Fig. 3 Global palaeogeography during the late Mississippian with location of occurrences of the Lonsdaleiinae. The geographical areas that are not labelled are listed in Fig. 4 (modified from Webb, 2002 and Somerville *et al.*, 2020).

of it in Ukraine, the Lviv-Volynsk Basin). The same problem exists with the Balkans, where there are some important studies (Kolosvary, 1951; Kostic-Podgorska, 1957, 1958), but the knowledge is scarce and incomplete. In other cases, the palaeogeographic domains do not fit completely with the recent political divisions. In order to avoid unnecessary terminology, we will use a generic geographic name as an approximation. Thus, Scotland, northern England, NW England, Derbyshire and North Wales can all be included in northern Britain, with South Wales, Bristol district and Mendips in southern Britain. Although two distinct palaeogeographic regions can be distinguished in Belgium, the Campine and the Franco-Belgian basins, we will consider both under the common name of Belgium, because the data in the Campine Basin are relatively scarce (Poty, 1981; Aretz and Chevalier, 2007) and because the two genera involved are respectively in the Franco-Belgian Basin and in the Campine Basin (Poty, 1981; Denayer et al., 2011). Finally, Turkey comprises several terranes that were clearly separated during the Mississippian; in at least four of them there are records of Mississippian corals. In the Istanbul-Zonguldak terrane, whose position in the Visean is not clear (Okay and Tüysüz, 1999), there is no record of Lonsdaleia or Actinocyathus and it will not be considered here as a separate unit. The Tauride block is related to the north border of Gondwana (Denayer, 2015).

If the oldest known occurrence of *Lonsdaleia* is that of *L. praenuntia* in the lower Visean from northwestern England, we can assume that the genus appeared in that area of the British Isles during the early Visean. The maximum development of *Lonsdaleia* occurred in the late Visean. The great transgression that took place in that period (Conil and Lys, 1977; Ramsbottom, 1979; Herbig, 1998) facilitated the distribution of the coral planulae transported by currents for large distances, because they easily found marine areas to attach themselves to. *Actinocyathus* first appeared also in the late Visean and quickly reached a large area of distribution.

We summarize the occurrences of both genera in Figs. 3 and 4. A general view of their distribution in the Palaeotethys is given in Fig. 3 and a more detailed one with data on each genus is presented in Fig. 4. Only data for the western Palaeotethys are included in the latter, because considering larger areas would impede a precise recognition of the details. Going from south to north and from west to east, the occurrences are as follows.

In the Tindouf Basin of southern Morocco (1), Lonsdaleia first occurs in the uppermost Visean, and reaches the maximum abundance and diversity during the Serpukhovian. Actinocyathus first occurs in the Serpukhovian, and is still recorded in the lower Bashkirian (Rodríguez et al., 2013a, 2013b). In the Béchar and Reggan areas of Algeria (2), the only record of Lonsdaleia is from the Serpukhovian and Actinocyathus hasn't been recorded there (Semenoff-Tian-Chansky, 1985). We included here the eastern Tafilalt (near Erfoud, eastern Morocco) where neither Lonsdaleia nor Actinocyathus have been recorded (Aretz et al., 2013). The same situation happens in Jerada (3) and



Fig. 4 Palaeogeography of the western Palaeotethys with location of occurrences of the Lonsdaleiinae in the Visean, Serpukhovian and Bashkirian (modified from Webb, 2002 and Somerville *et al.*, 2020). 1 – Tindouf Basin; 2 – Béchar-Reggan; 3 – Jerada; 4 – Azrou-Khenifra Basin; 5 – Ossa-Morena; 6 – Betic Cordillera; 7 – Cantabrian Mountains; 8 – Pyrenean Mountains; 9 – Montagne Noire; 10 – Carnic Alps (+Nötsch); 11 – Nova Scotia; 12 – Ireland; 13 – Northern Britain; 14 – Southern Britain; 15 – Southern Belgium; 16 – Germany; 17 – Poland; 18 – Donetz Basin; 19 – Moscow Basin; 20 – Balkans; 21 – Tauride.

the Azrou-Khenifra Basin (4) of northeastern and north central Morocco, respectively (Said *et al.*, 2007, 2013; Aretz, 2010). In the Guadiato area and the Los Santos de Maimona Basin (Ossa Morena, SW Spain) (5) both genera are also absent in the Visean and Serpukhovian rocks (Rodríguez and Falces, 1994; Rodríguez *et al.*, 2016). On the contrary, in the Betic Cordillera (S. Spain, 6), *Lonsdaleia* and *Actinocyathus* were recorded in the uppermost Visean and lowermost Serpukhovian (Herbig and Mamet, 1985).

In the Cantabrian Mountains of North Spain (7) both genera are absent in the Visean, but they have been recorded in the Serpukhovian (Boll, 1985). In the Pyrenean Mountains (8), Perret and Semenoff-Tian-Chansky (1971) recorded *L. duplicata*, and in the Hautes Corbières, a massif located somewhat further north, Semenoff-Tian-Chansky and Ovtracht (1965) recorded *L. corbariensis*, in both cases in the upper Visean. *Actinocyathus* was not recorded in that region. In the Montagne Noire of south central France (9), *Lonsdaleia* was recorded both in the upper Visean and in the Serpukhovian. On the contrary, *Actinocyathus* was only recorded in the Serpukhovian (Aretz and Herbig, 2003; Poty *et al.*, 2019). In addition, Poty and Hecker (2003) described there the subgenus *L. Serraphyllum*, with intermediate features between *Lonsdaleia* and *Actinocyathus* (Fig. 2). In the southern Austrian outcrops (Greywacke zone, Nötsch and Carnic Alps) (10), only the fasciculate genus *Lonsdaleia* has been recorded, both in the upper Visean from the Greywacke zone (Heritsch, 1933) and in the Serpukhovian of the Carnic Alps (Rodríguez *et al.*, 2019).

In Nova Scotia (11), Lonsdaleia has been also recorded in the upper Visean, but not in the Serpukhovian (Lewis, 1935; Poty, 2002). In Ireland (12), both genera, Lonsdaleia and Actinocyathus have been recorded (Caldwell and Charlesworth, 1962; Gallagher and Somerville, 1997; Somerville et al., 2007), including transitional forms between them (Cózar and Somerville, 2005). A similar occurrence is recorded in the north of England and Scotland (northern Britain, 13), where both genera occur in the upper Visean, but only Actinocyathus has been verified in the Serpukhovian (Hill, 1938-41; Jackson, 1958; Mitchell, 1989; Riley, 1995). In southern Britain (14), the record is like that in Ireland, with the two genera occurring both in the upper Visean and Serpukhovian (Mitchell, 1989; Riley 1995). In southern Belgium (15) the detailed studies by Poty (1985, 1989) indicate the presence of Actinocyathus in the coral zone RC8, i.e., uppermost Visean. Poty et al. (2006) showed also the presence of Lonsdaleia in that zone. As stated previously, the record in Germany is scarce and derived mainly from allochthonous rocks. Actinocyathus has only been recorded in the Brigantian from the Rhenohercynian domain (16), and there is no record of Lonsdaleia. Similarly, only Actinocyathus has been recorded in the upper Visean from the Lublin Basin (Poland) (Khoa, 1977) (17) and its southeastern prolongation in the Lviv-Volynsk Basin (Ukraine) (Shul'ga and Ogar, 2009). The record in the Serpukhovian from Poland is very poor and composed mainly of undissepimented corals, and neither Lonsdaleia nor Actinocyathus have been reported there. Both Lonsdaleia and Actinocyathus have been recorded in the Donetz Basin (18) (Vassiljuk, 1960; Fedorowski, 2022). In the Moscow Basin (19), the Lonsdaleiinae flourished, with Lonsdaleia and Actinocyathus, both being present in the upper Visean and in the Serpukhovian (Dobrolyubova, 1958; Hecker, 1997, 2001, 2010; Somerville et al., 2020).

None of the Lonsdaleiinae genera have been cited in the Balkans (20). As stated previously, the record there is scarce, and that region seems to be a mélange of terranes that were piled up against the Variscan Cordillera during the orogeny. The Tauride (21) represent a terrane closely related with northern Gondwana.

The Urals and Timan provided occurrences of *Lonsdaleia* and *Actinocyathus*, both in the upper Visean and Serpukhovian (Gorsky, 1949; Degtiarev, 1965; Sayutina, 1973; Kossovaya, 1996, 1997). Gorsky (1935, 1938) described several species of *Lonsdaleia* and *Actinocyathus* in the upper Visean from Novaya Zemlya, but only the latter in the Serpukhovian.

China and the surrounding areas have usually been divided into four or five areas regarded as different terranes with different sedimentation and coral assemblages during the Carboniferous (Wu and Zhao, 1979; Yang et al., 1985; Wang, 1989): In Tianshan Mountains, Wang et al. (1994) cited Lonsdaleia in the lower Visean. It is a very important recording, because the only other mention to lower Visean Lonsdaleia is that of *L. praenuntia* in northern England. In the same region, Lonsdaleia has been recorded in the upper Visean (more precisely in the Jilin area, Yang et al., 1985). In Kunlun Mountains, northwestern China, Wu et al. (1982) described Lonsdaleia in the upper Visean. Lonsdaleia and Actinocyathus have been mentioned and/or described in the upper Visean and Serpukhovian from Sichuan, Guizhou, and Hunan provinces of South China (Yu, 1937; Jia, 1977; Fan, 1978; Jiang, 1982; Wang et al., 2006), but there are no mentions of these genera in North China.

Fontaine *et al.* (1991, 2003, 2005) cited or described *Lonsdaleia* and *Actinocyathus* in the upper Visean and Serpukhovian from Thailand and Malaysia.

In Japan, there are several citations and descriptions of both *Lonsdaleia* and *Actinocyathus* in the upper Visean and Serpukhovian from different areas, such as Akiyoshi terrane, Ichinotani Mountains, Fukuji, Onimaru and Omi Limestone (Kato, 1966; Niikawa, 1979; Kato *et al.*, 1987; Igo and Adachi, 2001).

The only citing in Australia of *Lonsdaleia* (Hill and Woods, 1964) seems to be a misinterpretation and consequently the Lonsdaleiinae are not present on that continent.

In North America, Webb (1987) described both *Lonsdaleia* and *Actinocyathus* from the upper Chesterian (Serpukhovian) in Oklahoma and Arkansas. Fedorowski *et al.* (2012) described *Lonsdaleia* in the Bashkirian from Sverdrup Basin (Arctic Canada).

2.3. Insights for the evolution of the subfamily Lonsdaleiinae from the western palaeotethys

Lonsdaleia and Actinocyathus are present in all continents except in Australia and South America during the Visean and Serpukhovian. So, they can be considered as cosmopolitan genera. The configuration of their occurrences can be used for illustrating the coral migration patterns in the late Mississippian. The major factors controlling those patterns are.

- 1) The coral reproductive cycle; being sessile epibenthic organisms, their only way for migration and colonization of new habitats is the transport of the larvae (planulae).
- 2) The oceanic regime of currents; applying what we know on the system of currents in our planet, Earth's rotation causes a main equatorial current that flows from east to west, and other currents flow polarwards towards high latitudes conditioned by the distribution of the land masses. We can apply actualism here, because the Earth's rotation did not change during its history.
- 3) The distribution of land masses; we have an approximate view of the position of continents because of the many studies that have been carried out from different perspectives (tectonic, palaeomagnetic, biogeographic, etc.). But still, there are different models and reconstructions that do not agree on this matter (see Golonka, 2002; Blakey, 2008; Metcalfe, 2013; Scotese, 2021). The main reason is that the Visean and Serpukhovian are synorogenic periods and the land masses were changing at that time.

In order to identify the migration paths of the studied genera, the first step is determining where they appeared. The main distribution of *Lonsdaleia* and *Actinocyathus* occurred during the late Visean, but there are at least two earlier records. The two places where *Lonsdaleia* has been mentioned in the early Visean are NW England and Tianshan Mountains (NW China). It is not possible that it migrated between the Tianshan terrane and the UK during the early Visean without leaving occurrences in geographically intermediate areas, many of which are quite well studied. This implies that either one or both of these references may not be correct or that the genus could be polyphyletic.

Consequently, three hypotheses can be proposed on the origin of the genus.

1. Lonsdaleia praenuntia appeared in Arnside, NW England, near the western end of the Palaeotethys in the early Visean. It had a certain period without main changes and without expanding its distribution. But in the late Visean a major transgression took place and Lonsdaleia developed some structural changes diverging in several species in the same geographical area (L. alstonensis, L. caledonia, L. melmerbiensis, L. siblyi, etc.). It also evolved into the massive forms of the genus Actinocyathus (Smith, 1915). Both genera, also favoured by the transgression, migrated quickly. Theoretically, the main current in the Rheic Ocean, which was near to its final closure, should flow south-westwards (Fig. 4). Thus, the planulae could reach quickly the Nova Scotia territory (Poty, 2002). This hypothesis is also supported by foraminiferal data (cf. Cózar and Somerville, 2021). In a different way, the planulae could use local currents in the shallow platforms that bordered the continent of Laurentia plus Baltica towards Scotland, southern England, Belgium, Poland, Donetz and the Russian Platform. The migration was very guick, and both genera reached the Urals, the Tianshan Mountains, North and South China, and Japan, along the northern margin of the Palaeotethys (Fig. 3). The maximum area of development was the Russian Platform (Hecker, 1997, 2010; Poty and Hecker, 2003). A notable diversification took place there, with the development of two evolutionary lines in Actinocyathus and the appearance of the species group of L. corbariensis in Lonsdaleia, adapted to less favourable environments, and related to microbial communities. That line is represented in the Russian Platform by L. tichyi.

In all the cited areas, *Lonsdaleia* and *Actinocyathus* flourished during the late Visean and Serpukhovian.

The local absence of them can be due more to problems of preservation and the destruction of the original habitats during the Variscan Orogeny. But their presence illustrates the quick evolutive and ecologic success of both genera in so many and distant geographic areas.

Also, during the late Visean, a migration took place towards the southeast, probably from the Moscow or Donetz Basin. The migration in this way was represented mainly by the L. corbariensis group of species, which has been recorded in the southern terranes of Europe, and mainly in the border of the Armorica-Iberian Massif (Montagne Noire, Pyrenees, Cantabrian Mountains and Betic Cordillera). Most occurrences of these species are related to allochthonous blocks in flyschoid facies or in olistostromes and related to microbial communities (Boll, 1985; Herbig and Mamet, 1985; Aretz and Herbig, 2003; Poty and Hecker, 2003; Rodríguez et al., 2019). But the migration towards southwest reached longer, far away to the epicontinental platforms in North Africa (Legrand-Blain et al., 1989). There, the represented species group of Lonsdaleia is the L. duplicata group that reached the region during the late Visean and Actinocyathus, which reached the region during the Serpukhovian (Rodríguez et al., 2013b and above). The stable conditions in that region, far from the main tectonic movements occurring in the northern part of the western Palaeotethys, allowed some representatives of the genus to survive into the lower Bashkirian. A possible way of migration along the southern border of the Palaeotethys is discarded, because neither Lonsdaleia, nor Actinocvathus have been recorded in eastern terranes such as the Tauride Alborz etc. The reasons could be excessive amounts of terrigenous material or too high latitudes (see Denayer, 2015).

An additional route of migration took place from the Russian Platform towards the north, reaching Timan (Kossovaya, 1997) and Novaya Zemlya (Gorsky, 1935, 1938). This migration was directed to the west and reached the Sverdrup Basin in northern Canada (Fig. 3), where *Lonsdaleia* has been recorded in the Bashkirian (Fedorowski and Bamber, 2012; Fedorowski *et al.*, 2012).

The occurrence of *Lonsdaleia* and *Actinocyathus* in the Midcontinent and Western Interior of USA (Webb, 1987) is less easy to explain. The northern route is discarded because these genera have not been recorded in the upper Visean and Serpukhovian in the Rockies. The southern route along the Rheic Ocean is also difficult to accept, because most palaeogeographic reconstructions close the Rheic Ocean before the Visean (Blakey, 2008; Cao *et al.*, 2017; Scotese, 2021). However, the coral data (as well as

foraminiferal data; see Cózar and Somerville, 2021) show that the Rheic Ocean didn't close before the Bashkirian, because there are typical components of the western Palaeotethys assemblages in the Midcontinent and Western Interior domains (García-Bellido and Rodríguez, 2005; Rodriguez and Kopaska-Merkel, 2014). However, the percentage of western Palaeotethys species in North America is very low, indicating a low level of communication.

A particular case is that of the domains of southwestern Spain, the Jerada Basin and the Azrou-Khenifra Basin, where *Lonsdaleia* has not been recorded (localities 3–5, Fig. 4), despite the fact that there have been intense studies on the coral assemblages (Aretz, 2010; Said *et al.*, 2013; Rodríguez *et al.*, 2016). The explanation may be that both the Iberian Massif and the Anti-Atlas formed important barriers for this genus. On the other hand, these domains could be separated from Nova Scotia and Western European domains by a narrow but a deep remnant of the Rheic Ocean, before its final closing. It is remarkable, because in most other aspects, the assemblages of these areas have a high level of similarity with other domains from North Africa and Western Europe.

- 2. The second hypothesis only changes the origin of the genus to the northern Tianshan Mountains, if Lonsdaleia praenuntia is really an Axophyllum. From this area, it migrated to the east and southeast to the North China, South China and Japan terranes and at the same time to the west along the coastal platforms of the terranes bordering the northern Palaeotethys, and/or along the equator, taken advantage of the equatorial current and possible islands in the Palaeotethys, reaching first the Eastern European Platform and later (always in the late Visean) the Western European region and the northern Gondwana epicontinental seas. This second hypothesis has the advantage that it could explain better the late arrival to the latter and the scarcity of occurrences in the Midcontinent and Western Interior from North America. It also would explain better the absence in the basins of the South Iberian Massif and in the most northern basins of Gondwana and the refuge in Sverdrup, one of the most distant domains in which Lonsdaleia has been recorded.
- 3. The appearance of *Lonsdaleia* and *Actinocyathus* almost simultaneously in the whole Palaeotethys during the late Visean and two doubtful occurrences in the early Visean allow the validity of a third hypothesis: these genera could be polyphyletic, and originating in different regions at the same time (South China, northern Tianshan Mountains of northwestern China, Russian Platform, and

northern Britain). The homogeneity of structures and microstructures reduces the possibilities of this last hypothesis, but it is not completely discarded. The appearance in different domains could be favoured by the creation of conducive environments for colonial corals associated with the Asbian transgression, which covered with epicontinental seas many previous coastal plains. It also explains the quick diversification. All the routes described in the first hypothesis are valid also in this one.

The distribution of the occurrences of Lonsdaleia and Actinocyathus in the western Palaeotethys (Fig. 4) confirm the proposal of Somerville et al. (2013) on their division into four palaeogeographical subprovinces that we augment to six: (1) The Atlantic Subprovince (West European countries and Nova Scotia) shows a consistent abundance of both genera in the Visean. In some areas of this subprovince, the Variscan Orogeny activity implied increase of terrigenous sediments and reduction of coral habitats, with the subsequent extinction of these genera during the Serpukhovian. (2) The Mediterranean Subprovince (Pyrenees, Montagne Noire, Betic cordillera, Rif, Balearic Islands, Nötsch, and Carnic Alps) shows similar occurrences of the L. corbariensis group in the upper Visean and/or the Serpukhovian. (3) The Saharan Subprovince (Béchar, Reggan, Ahnet-Mouydir, and Tindouf Basin) shows irregular occurrences of both genera, but it was a refuge for the genus Actinocyathus in the Bashkirian. (4) The West peri-Gondwanan Subprovince (SW Spain and Moroccan Meseta) shows a total absence of both genera. (5) The Eastern European Subprovince (Moscow Basin, Donetz Basin, and Voronezh) shows also abundance of both genera in the Visean and Serpukhovian. (6) Finally, the Central Europe Subprovince (Germany and Poland) shows an irregular record due to the effects of the Variscan Orogeny that produced a structural mélange of domains and occurrences, mainly in allochthonous facies.

3. Conclusions

The most likely ancestor of the genus *Lonsdaleia* is the solitary genus *Axophyllum*, which would acquire colonialism in favourable environments. *Actinocyathus* descends from *Lonsdaleia* by acquiring the massive habit by increasing the degree of integration in the colonies.

Two records of *Lonsdaleia* have been cited in the lower Visean from northern Britain and northern Tianshan Mountains of northwestern China. But none of these records have continuity in the middle Visean. The diversification and migration of *Lonsdaleia* and *Actinocyathus* only occurred in the late Visean and continued during the Serpukhovian. Three hypotheses are proposed on the origin and migration of these genera: (1) Origin in northern Britain and first migration along the epicontinental seas bordering the Palaeotethys; (2) Origin in northern Tianshan Mountains and first migration using the equatorial current and later the epicontinental seas; (3) Simultaneous appearance in several domains of the Palaeotethys and migrations using the previously mentioned routes (polyphyletic hypothesis). In all cases, the migration was quick, because these two genera are recorded in the upper Visean from the whole Palaeotethys and North America.

Lonsdaleia and Actinocyathus became extinguished in most regions at the end of the Mississippian. However, Lonsdaleia has been recorded in the Bashkirian of the Sverdrup Basin in northern Laurasia and Actinocyathus has been recorded in the Bashkirian of the Tindouf Basin in northern Africa. They both could represent refuges in tropical areas far from the zones affected by the Variscan Orogeny during the early Pennsylvanian.

The western Palaeotethys is divided into six subprovinces with different distribution of the Lonsdaleiinae and other rugose corals.

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Authors' contributions

All the authors participate in the planning of the study, in the review of the taxa, in the design of the manuscript, and wrote parts of it. All authors read the manuscript and approved it.

Conflict of interest

The authors declare that they have not conflict of interest.

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