

# **Space Weather:** Modelling the thermosphere ionosphere system during geomagnetic storms

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**Geospatial World Distinguished Scientist Forum**

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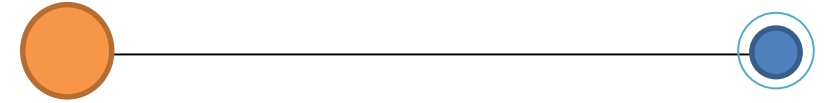
**12 April 2023**



Knowledge for Tomorrow



# Outline



## Space Weather

Upper Atmosphere

Geomagnetic storms

## Objective

Forecast the state of the upper atmosphere

## Models and Methods

Physics – based models : CTIpe

Data Assimilation (DA)

## Results

Impact of magnetospheric drivers in near-real time

Correcting the TI system state through data assimilation



# Space Weather



- **Space Weather** (Solar – Terrestrial physics) → **Solar activity**

- Sun and its atmosphere
  - Interplanetary space
  - Earth magnetosphere and upper atmosphere
- } **Geospace**

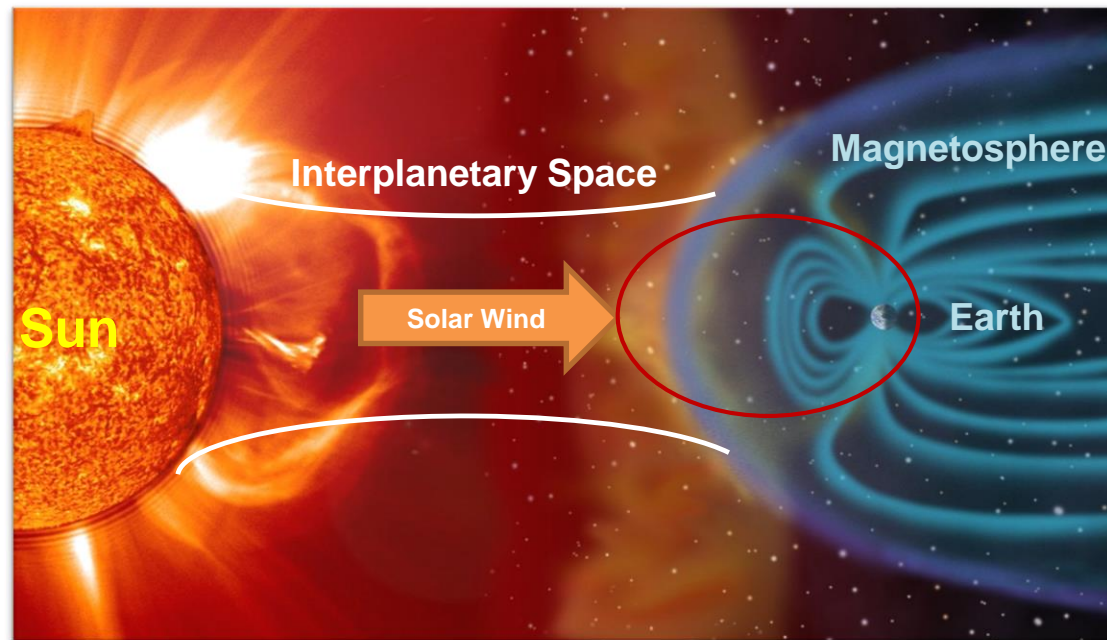
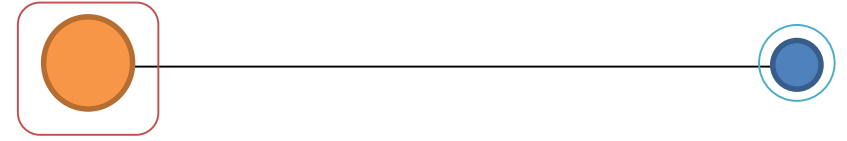


Image courtesy NASA



# Solar Flares and Coronal Mass Ejections



## Solar Flare

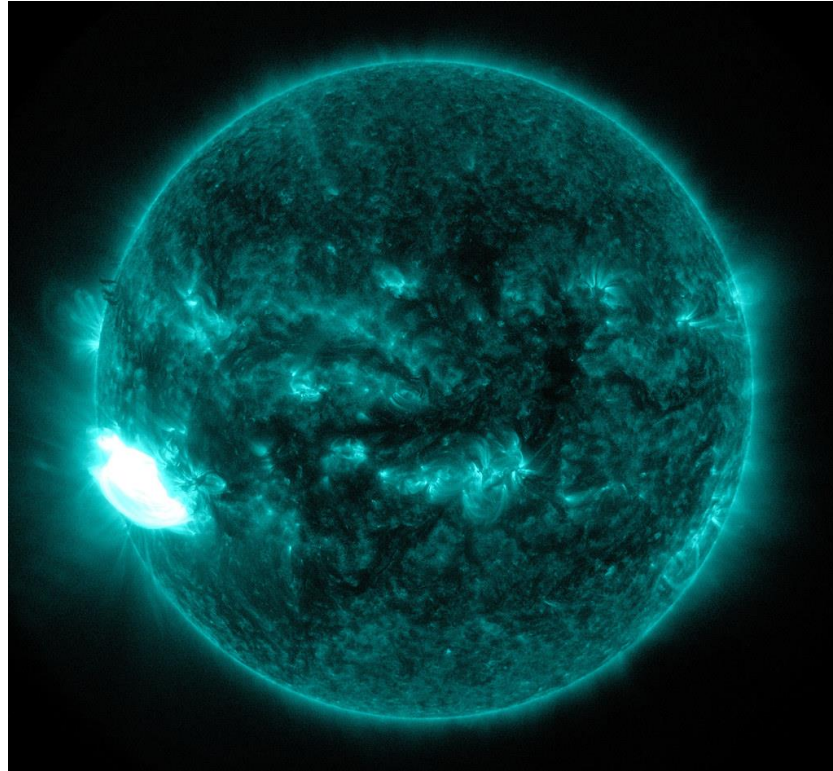


Image from NASA (SOHO mission)

## Coronal Mass Ejection

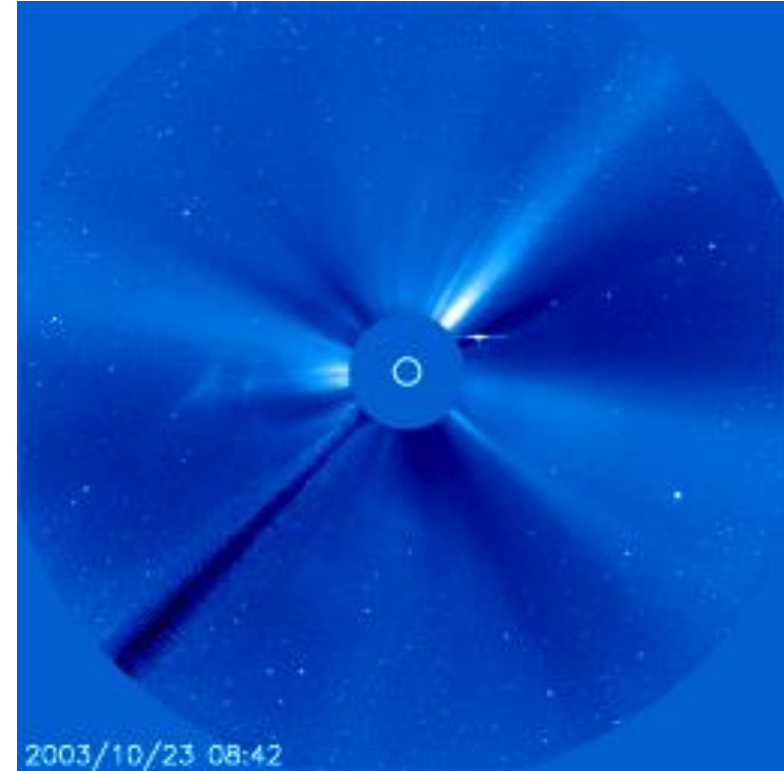


Image from NASA (SOHO mission)



# Interplanetary Space and Magnetosphere



## Interplanetary Space

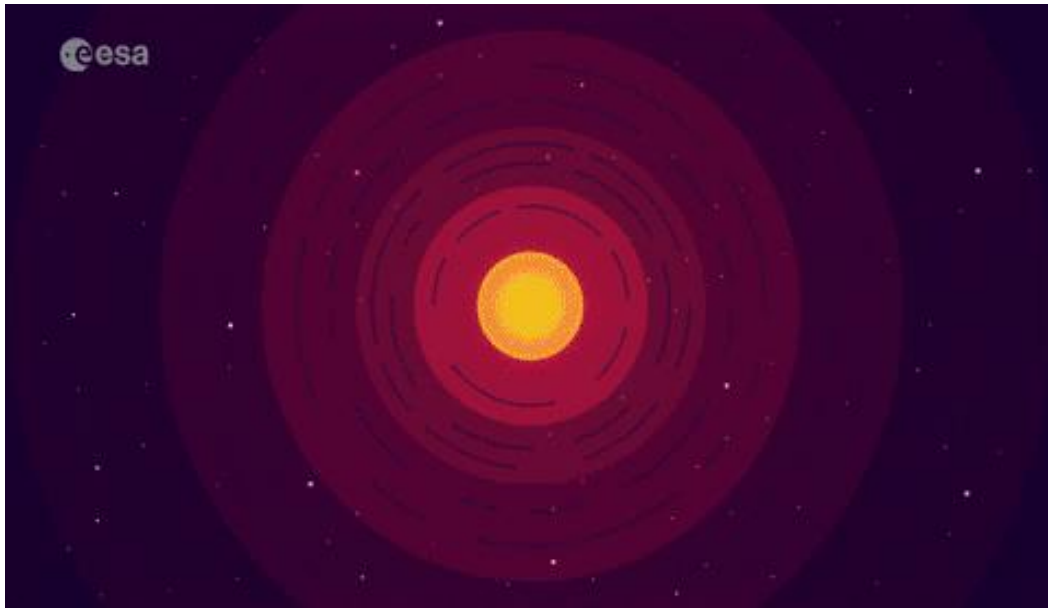


Image courtesy ESA

## Magnetosphere

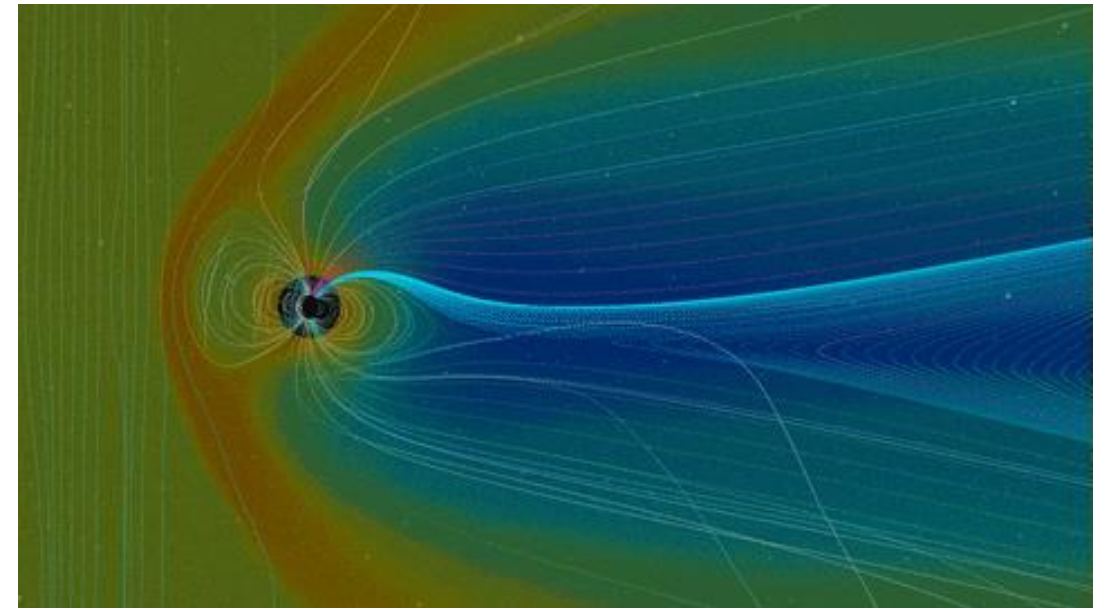
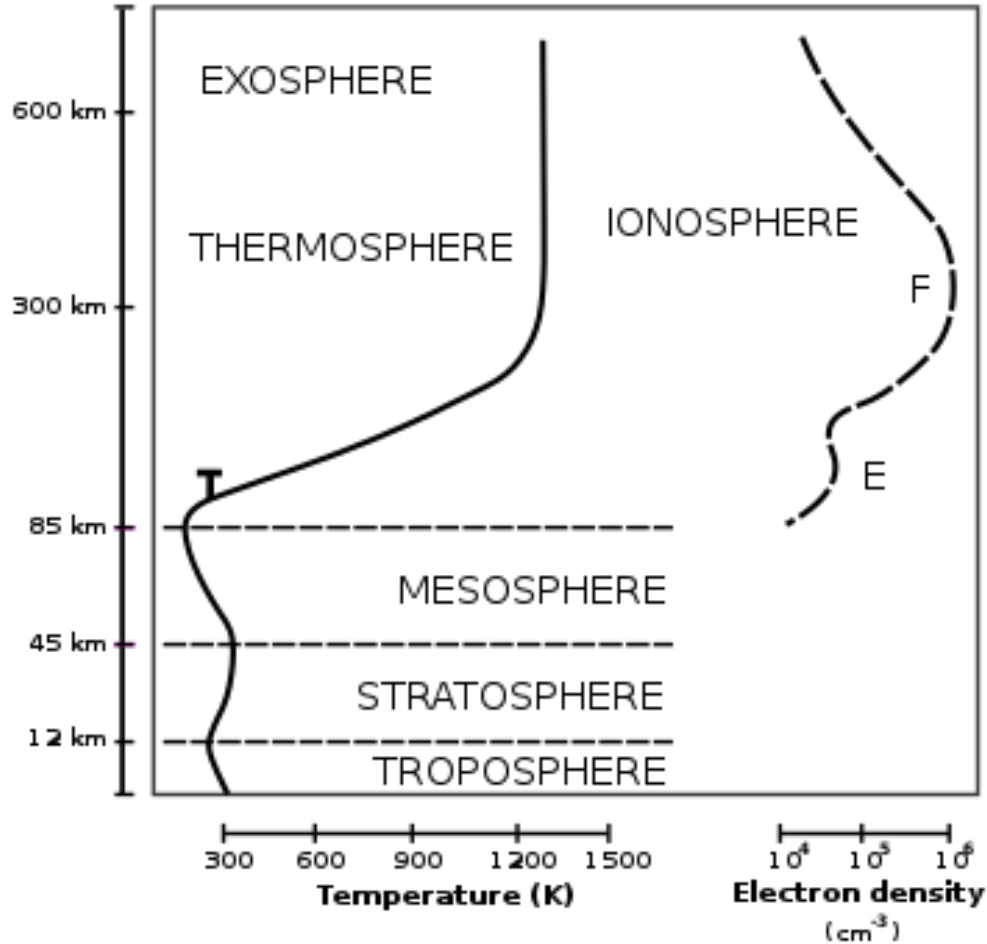
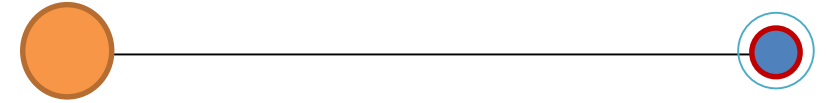


Image courtesy NASA



# Thermosphere – Ionosphere (TI) system

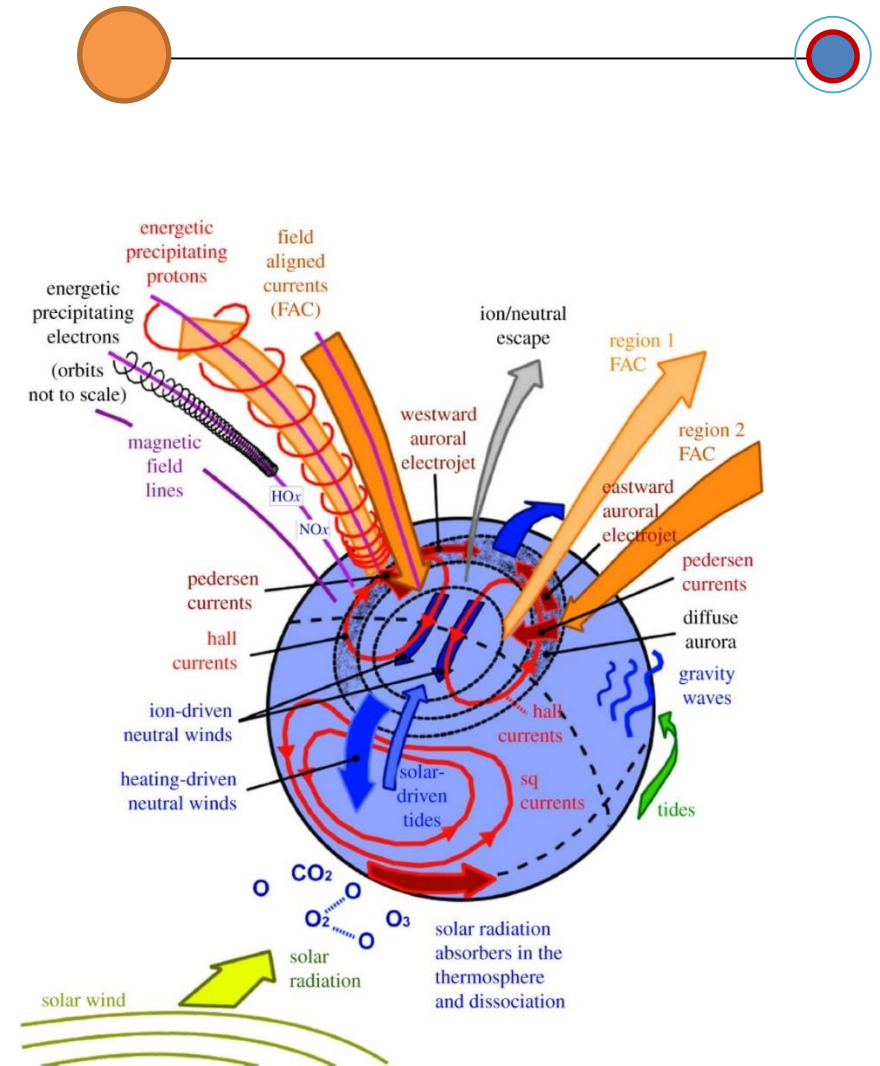


- **Thermosphere** (90,600)km: **Neutral composition**
  - Layer of the upper atmosphere between the mesosphere and exosphere.
  - Characterized by its **temperature** profile.
  - Increase of T with high at lower thermosphere and steady above 300 km
- **Ionosphere** (90, 2000)km: **Charged particles**
  - Characterized by high – densities of free charged particles – ions and electrons.
  - **Electron density** profile characterizes the changes of the ionosphere in altitude.
  - Coupling between transport heating, electrodynamics and chemistry.



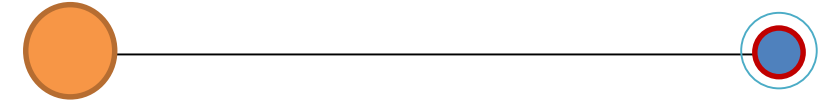
# Causes of TI variability

- TI system varies with altitude, latitude, longitude, time, season, solar cycle and **geomagnetic activity**.
- Main **driving mechanism**:
  - Solar ultraviolet (**UV**) and extreme ultraviolet (**EUV**) radiation.
  - Sun **energetic particles**.
  - **Solar wind** and **magnetic field**.
- TI processes and their interaction with the magnetosphere and the solar wind



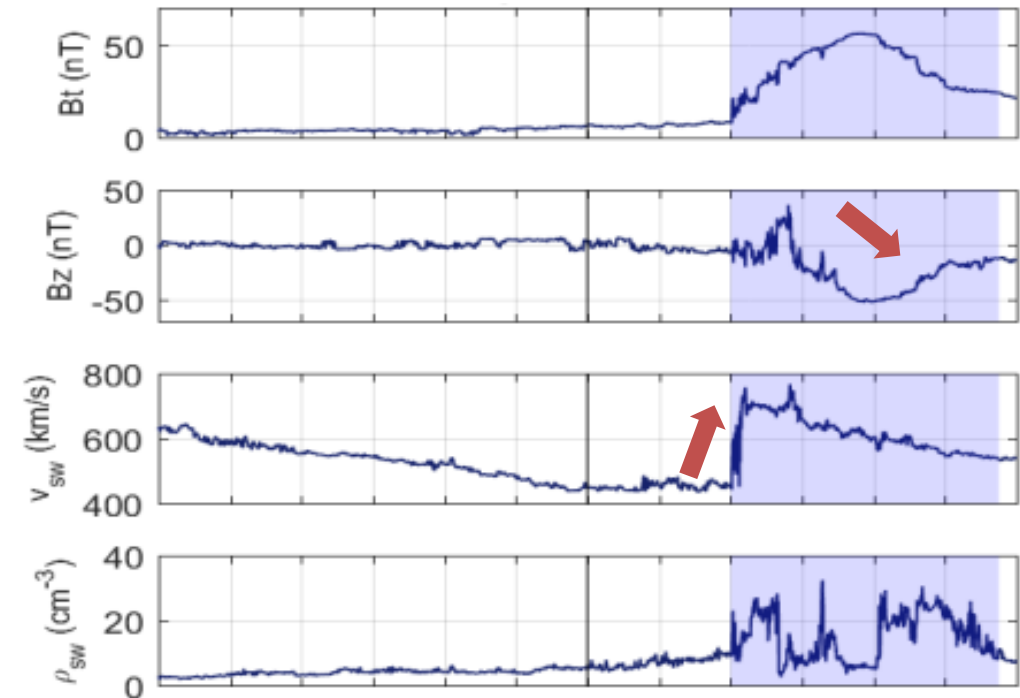
[Sarris, 2019]

# Geomagnetic storms



- **Geomagnetic storm:** Major disturbance of Earth's magnetosphere that occurs when there is an energy exchange from solar wind into the space environment surrounding Earth. Caused by solar flares or **coronal mass ejections** (CMEs)
- **Conditions:** High speed solar wind and **Southward directed solar wind magnetic field** ( $B_z$ ) opposing the Earth's magnetic field direction, at the dayside of the magnetosphere, sustained during several hours.

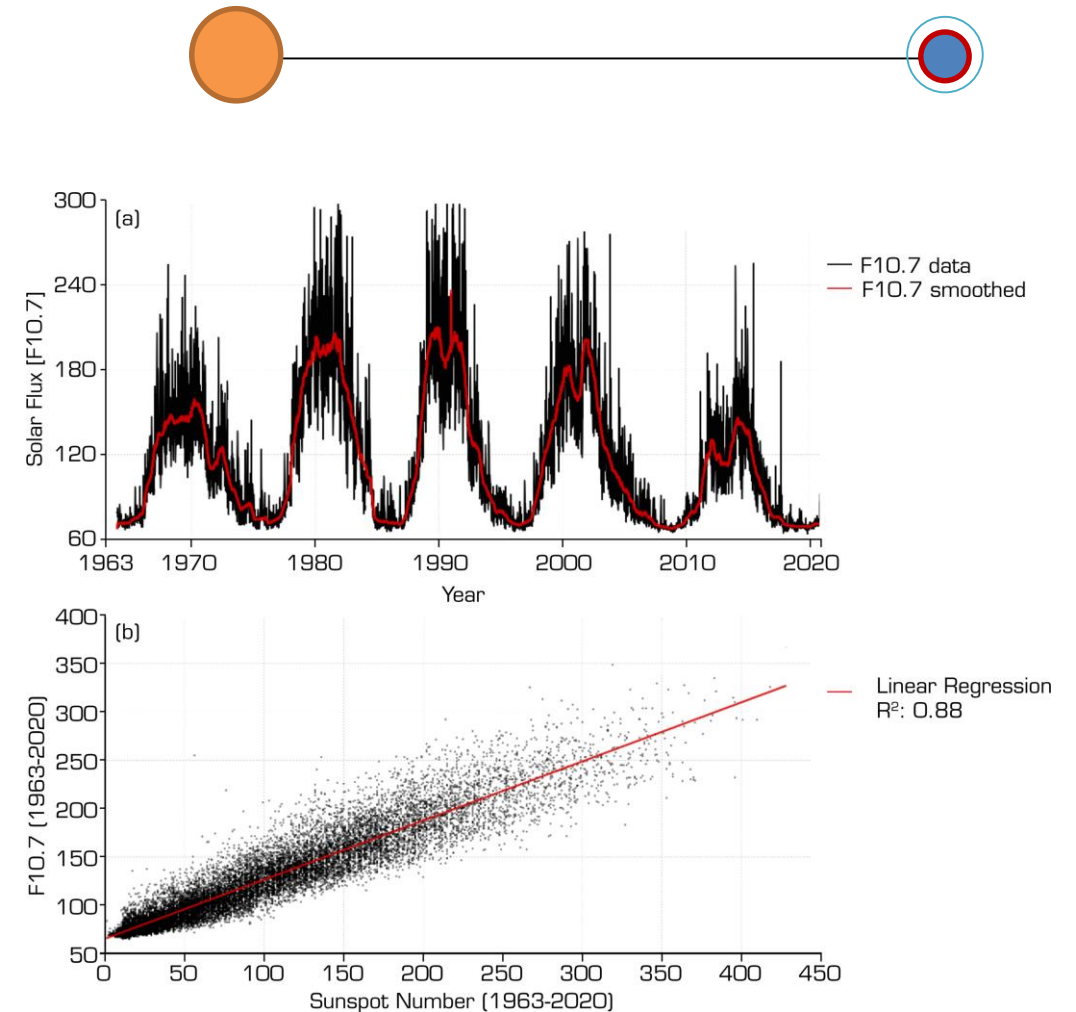
## Solar wind parameters





# Geomagnetic storms scales

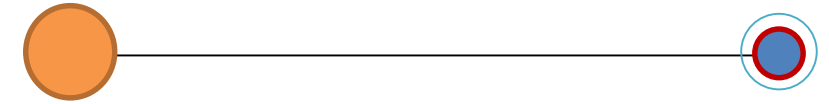
- **F10.7**: Solar flux at **10.7 cm radio emission** is a good indicator of solar activity. It correlates well with the sunspot number, UV and visible solar irradiance.
- **AE**: Geomagnetic index of the **auroral electrojet**, characterizes the maximum range of divergence from quiet geomagnetic level.
- **Kp**: Index to indicate the severity of the **global magnetic disturbance** of the near-Earth space, based on the range of variation of the magnetic field.
- **Dst**: GS result in intense currents in the magnetosphere. Equatorial current around Earth produces magnetic disturbances on the ground. A measure of this current is the **Disturbance storm time** index.



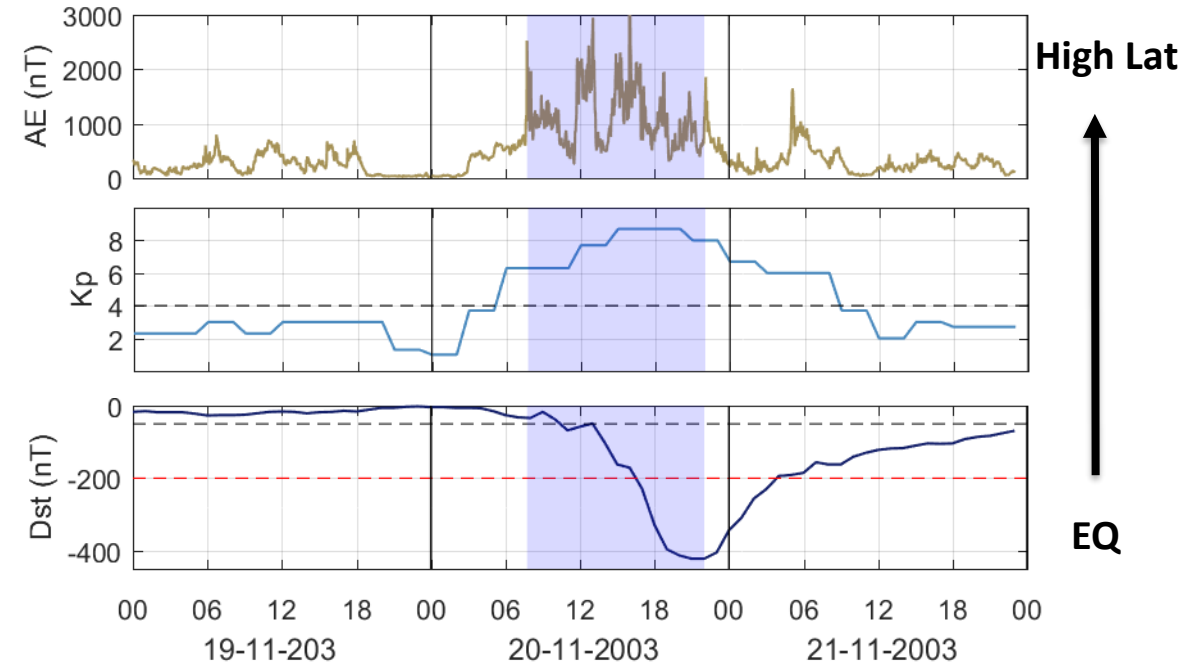
[Sousasantos, 2021]

# Geomagnetic storms scales

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## Geomagnetic conditions



# Why is so important the TI system?

- **Free charged particles affect propagation of electromagnetic signals.**
- **Communication** systems that utilize the ionosphere to reflect radio signals over long distances can be disrupted.
- Effects on **navigation**, like the Global Positioning System (GPS). Free electrons induce group delays and phase acceleration of the signal resulting in a bias in position determination.
- **Electric power** may be affected during geomagnetic storms. The presence of magnetic fields in the vicinity of a conductor (wire) induce electric currents.
- The fluctuation geomagnetic fields can induce currents into **pipelines**.

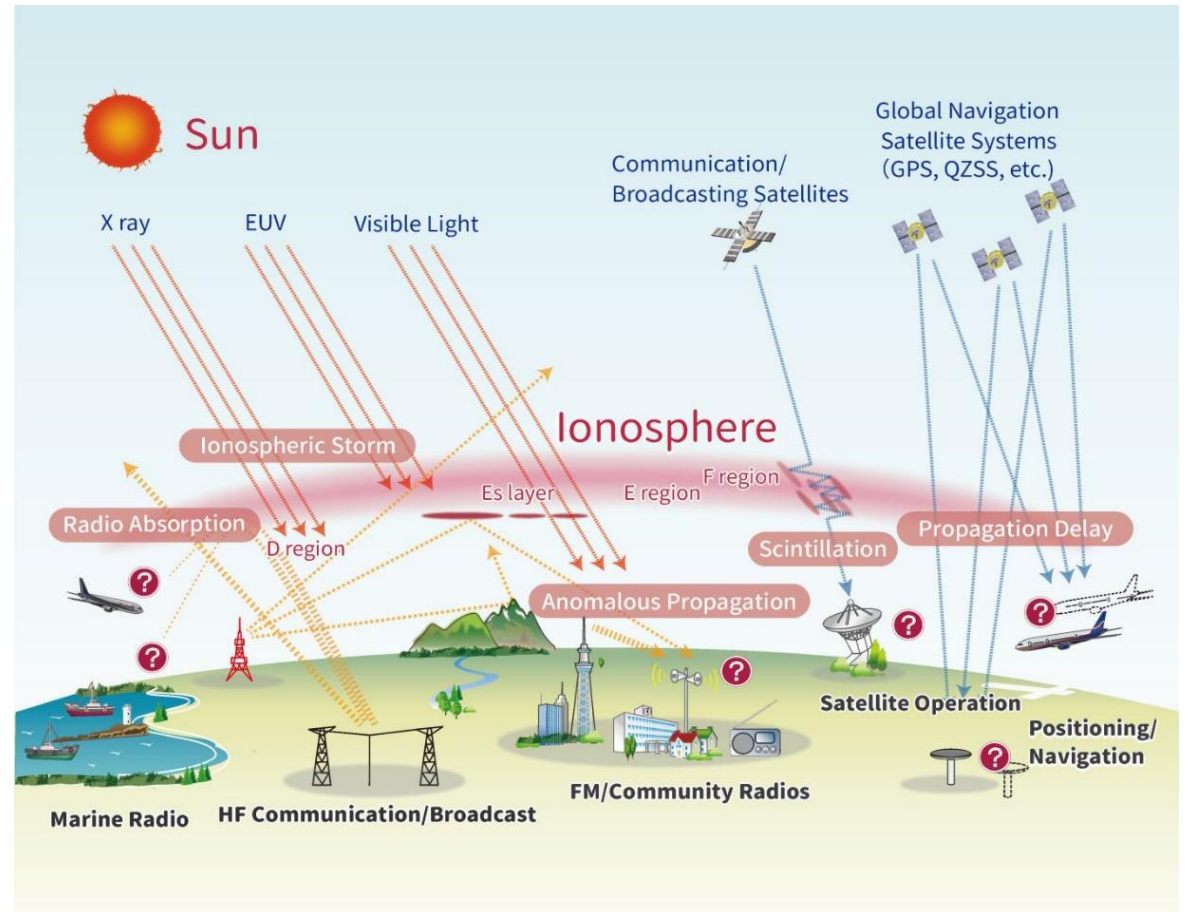
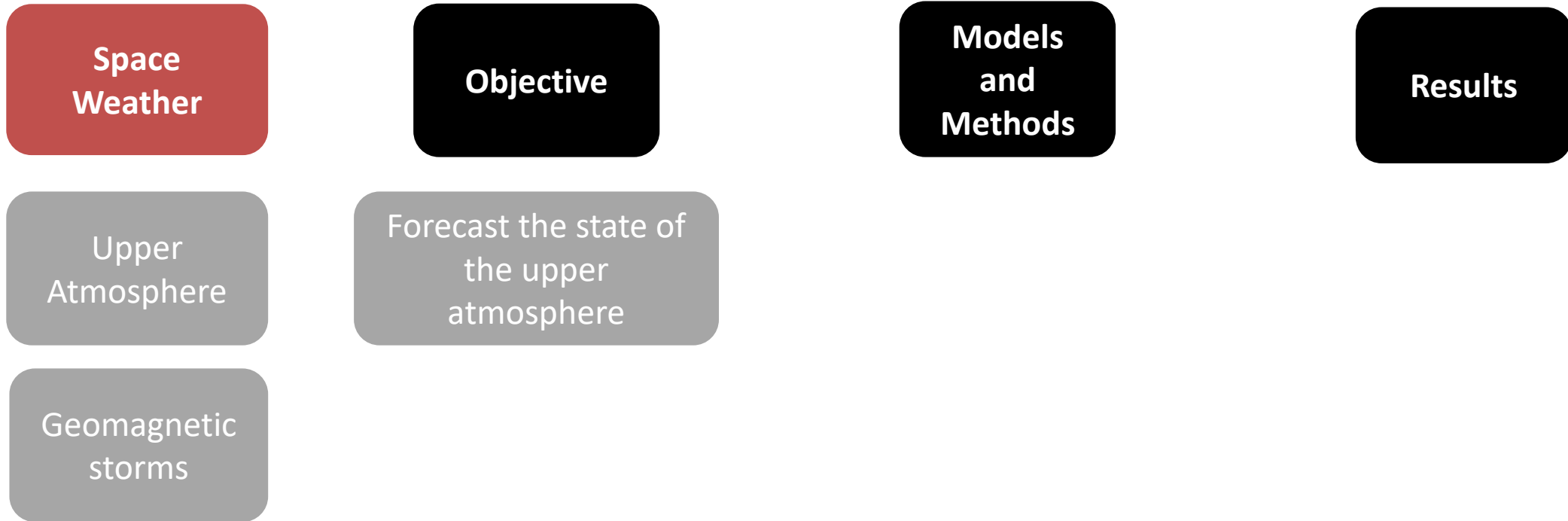
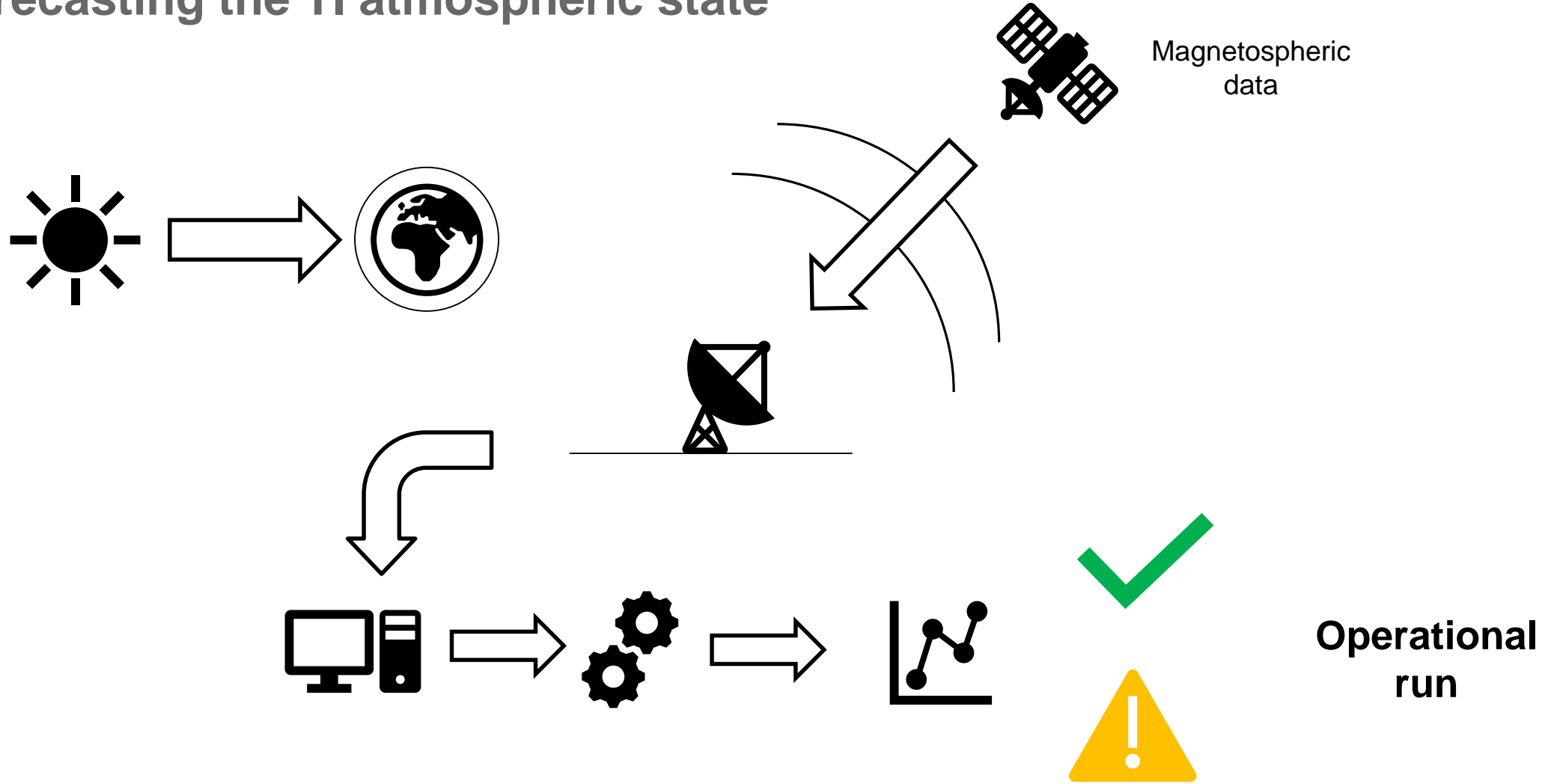


Image courtesy NIICT

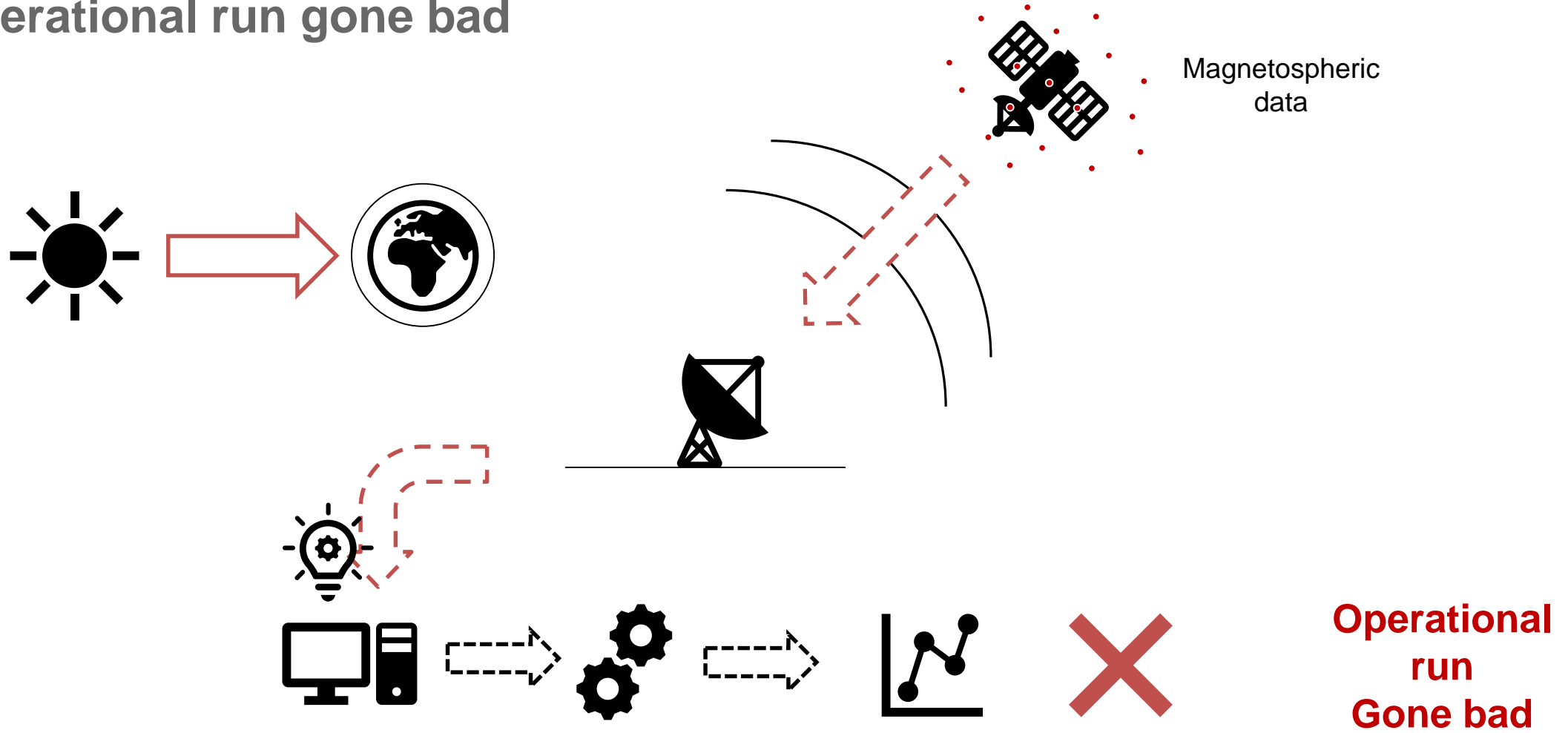
# Outline



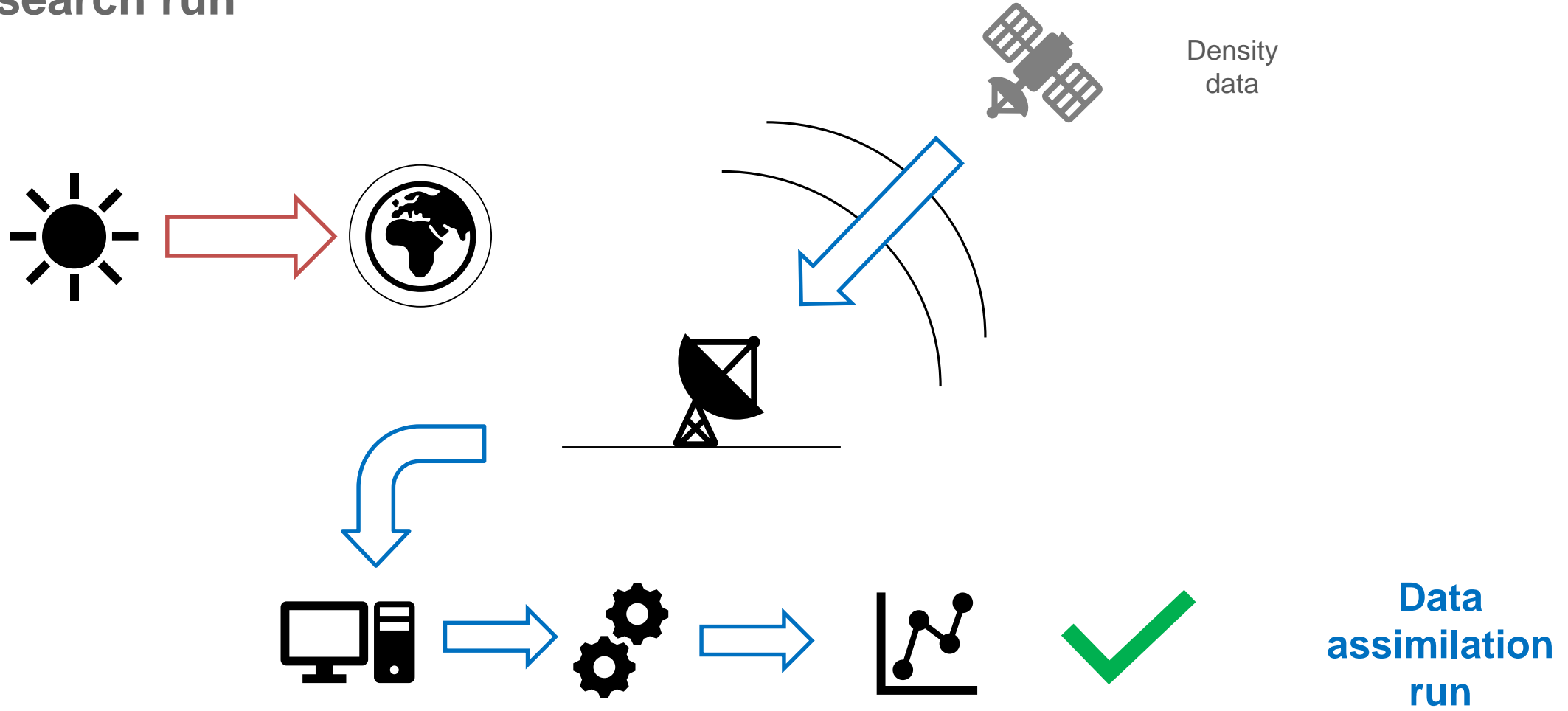
# Forecasting the TI atmospheric state



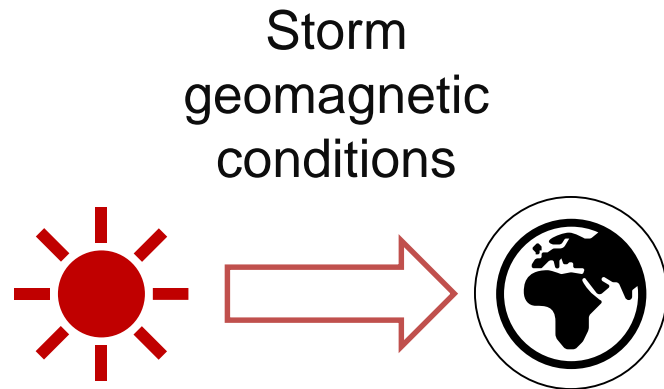
# Operational run gone bad




# Research run



# Objective: Operational vs. Research

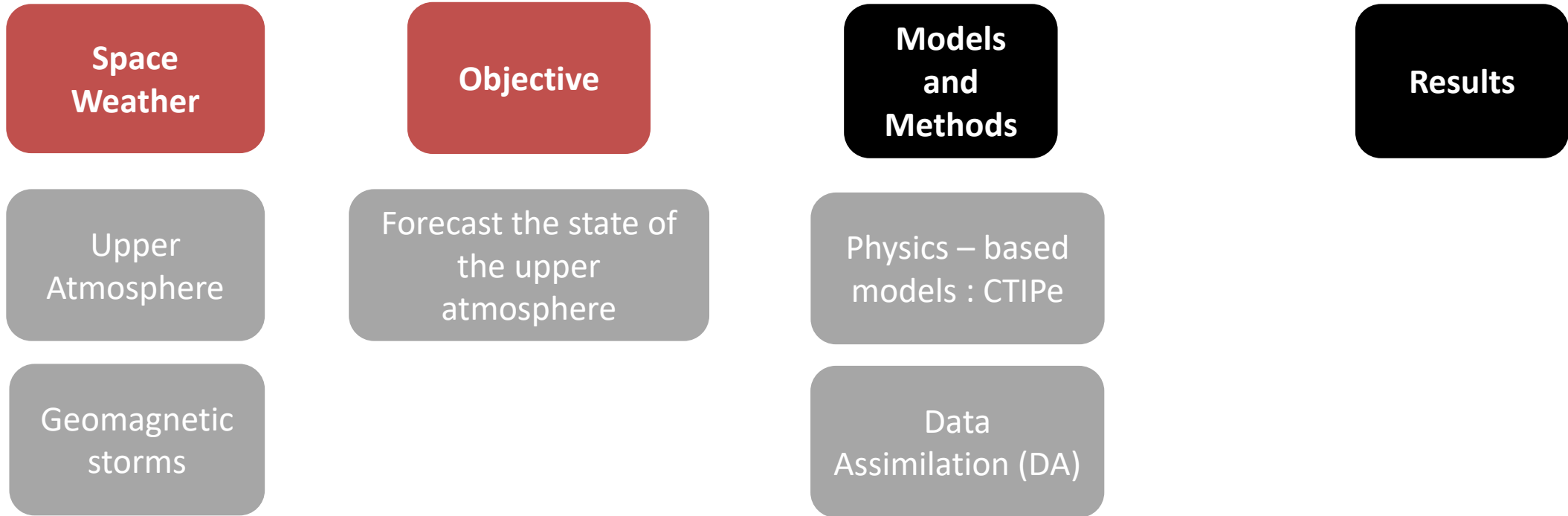


|  | Model run                | Input data                    | Study during storm conditions                             |
|--|--------------------------|-------------------------------|---|
| ①  | <b>Operational</b>       | Magnetospheric real time data | Impact of magnetospheric drivers in real time simulations |
| ②  | <b>Data Assimilation</b> | TI Observations               | Correcting the TI system state through data assimilation  |

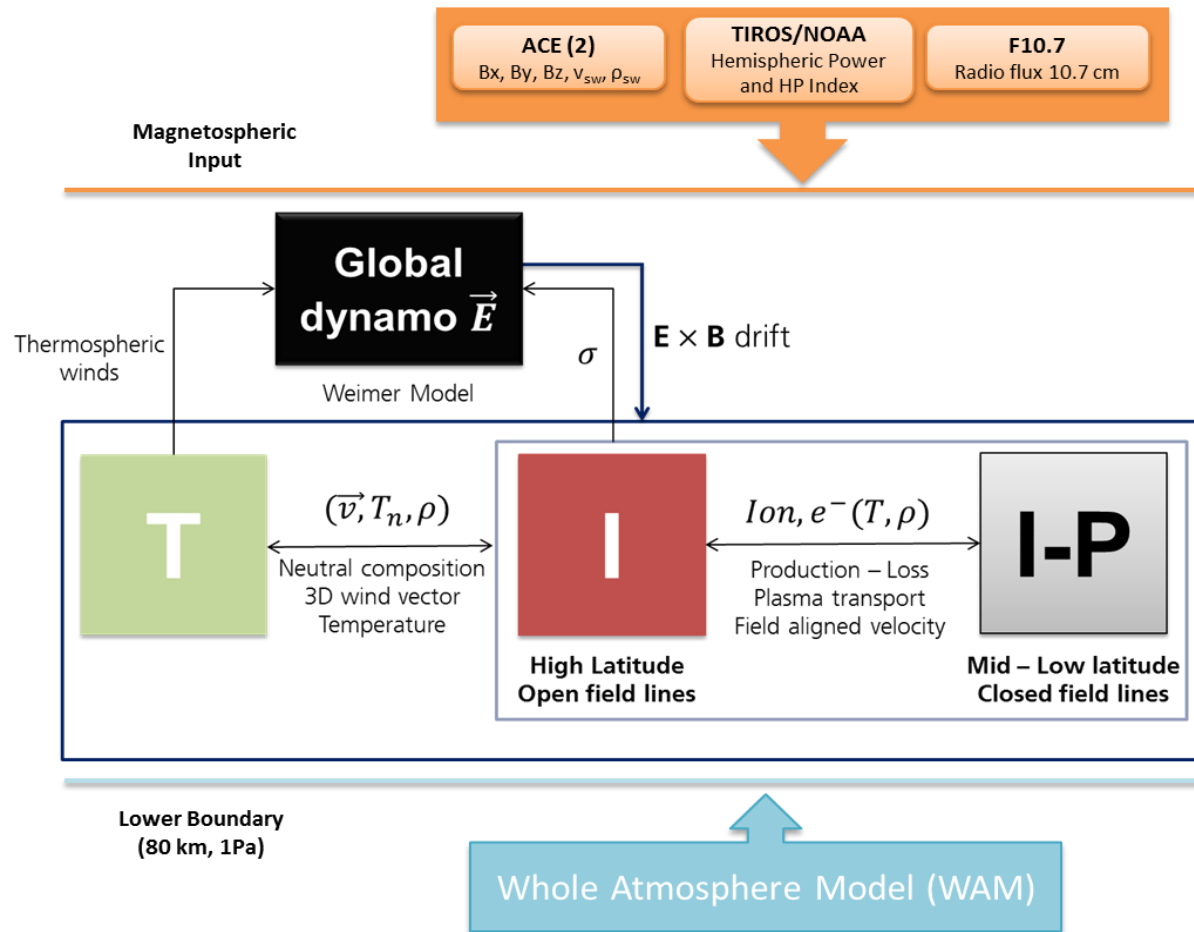




# Outline



# Physics – based model: CTIPe

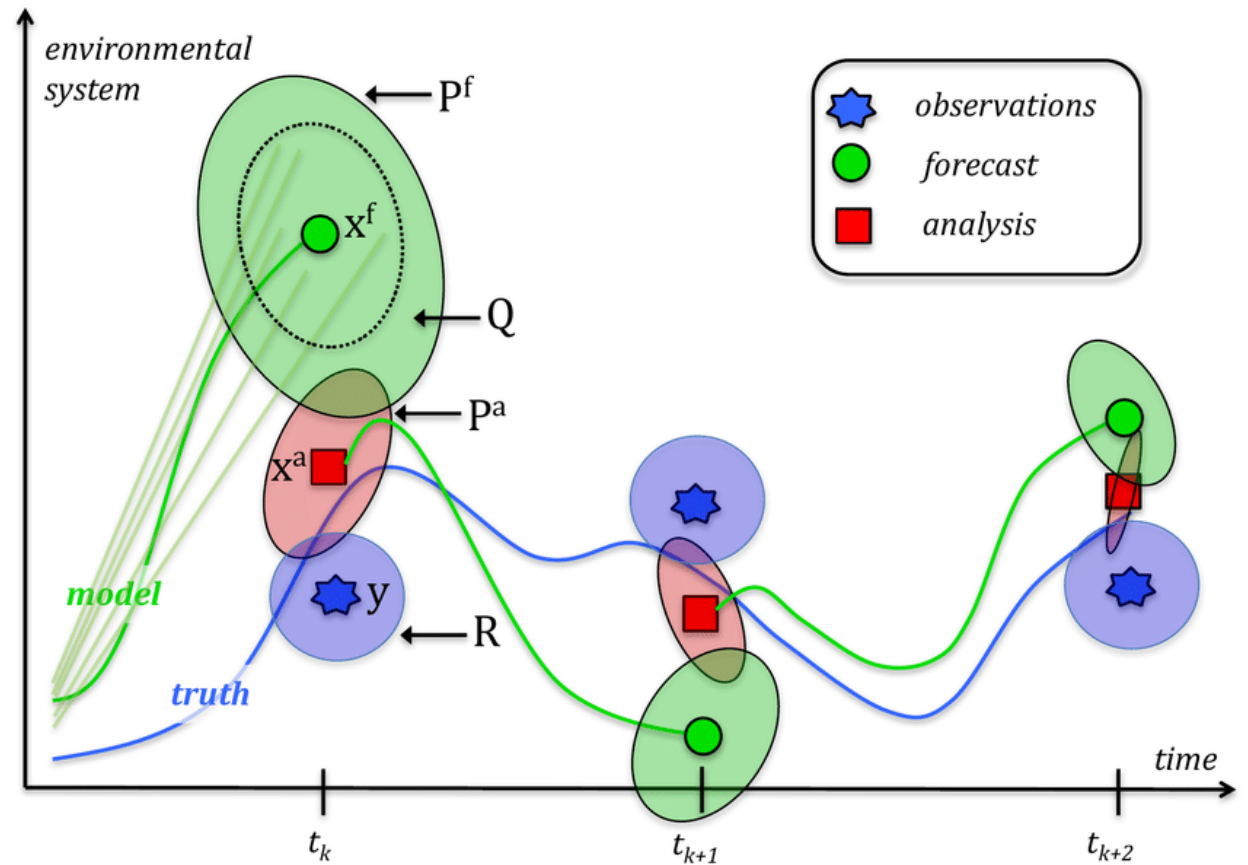


- Coupled Thermosphere Ionosphere Plasmasphere with Electrodynamics physics based model (SWPC-NOAA).
- Solves the equations of momentum, energy and composition for **neutral (T)** and **ionized (I, I-P) atmosphere**.
- **Global dynamo electric field:** Weimer Model
- **Magnetospheric input:** ACE measurements, TIROS/NOAA auroral precipitation and F10.7.
- **Lower boundary:** Simplified version of WAM at 80km



# Data Assimilation (DA)

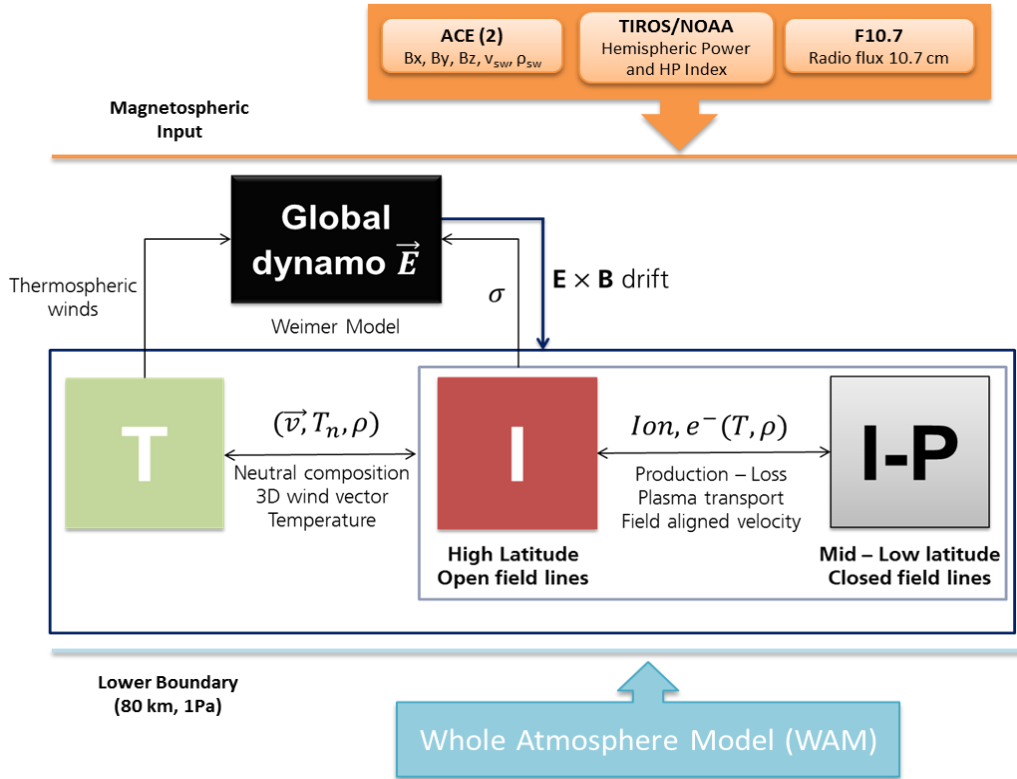
- **Data assimilation** combines different sources of information to estimate possible states of a system that evolves in time.
- Thermosphere – Ionosphere system is corrected by **combining density observations with models results** taking into account the **uncertainties** in each, while respecting certain constraints.



[Tandeo, 2018]

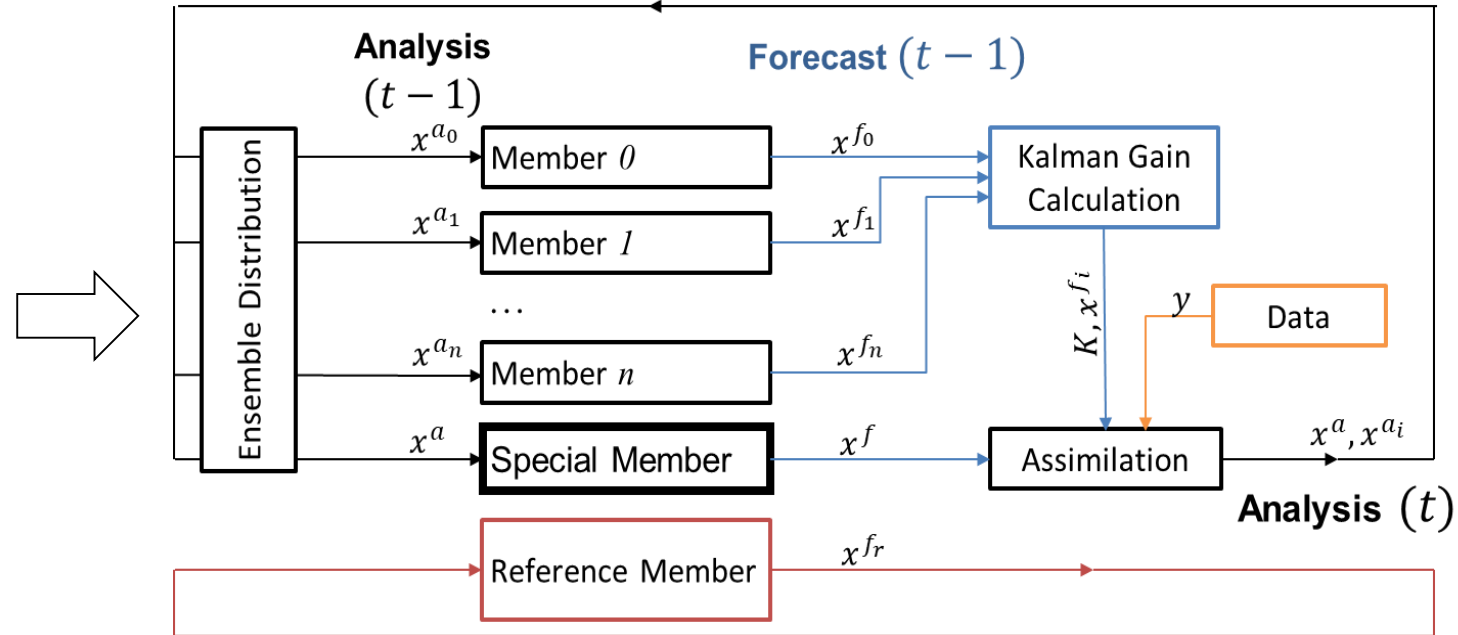
# CTIPe – TIDA: Physics based model with neutral mass density DA

## CTIPe



## TIDA

[S. Codrescu, 2018]

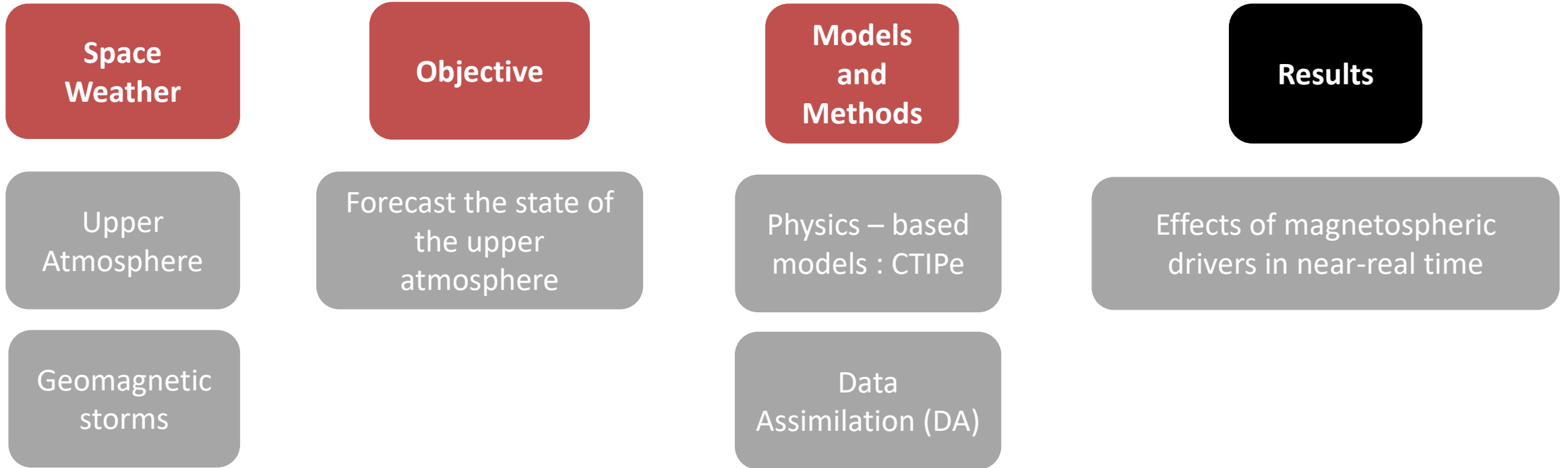


State Vector  $x = \begin{bmatrix} \text{model forcing} \\ \text{model state} \end{bmatrix} \longrightarrow \begin{matrix} x^a = x^f + K(y - h(x^f)) \\ y^f = h(x^f) \end{matrix}$  **KF update equation**

$$x = \{F_{10.7}, |v_{sw}|, \rho_{sw}, B_N, B_\theta, T_n, \gamma_O, \gamma_{O_2}, \gamma_{N_2}, M, U, V\}$$

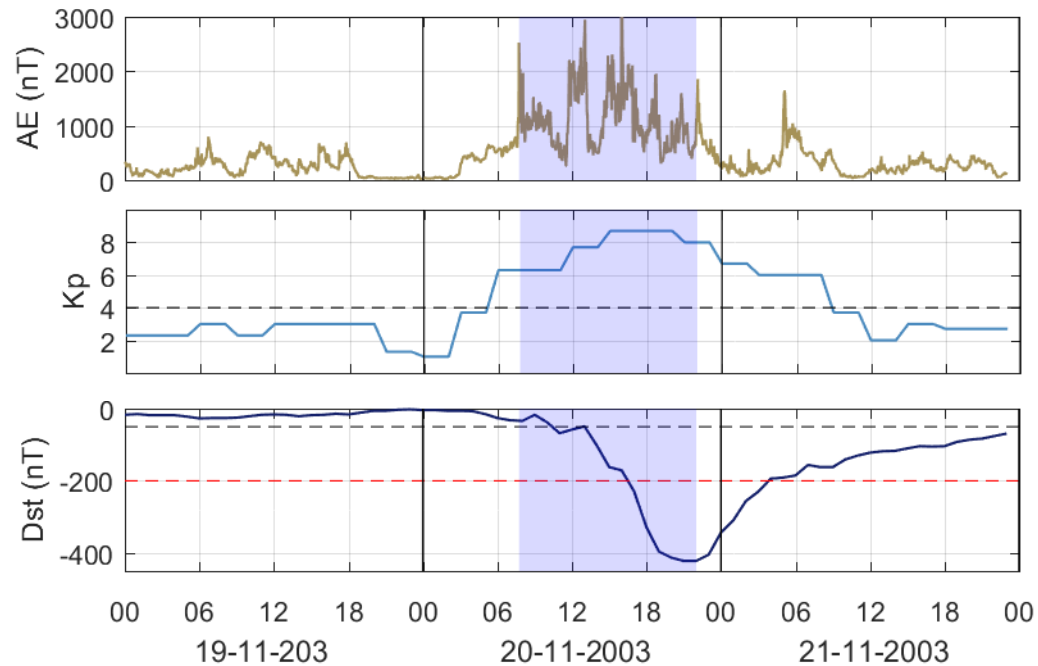


# Outline



# Case Study: 20<sup>th</sup> November 2003 Superstorm

## Geomagnetic conditions



High Lat

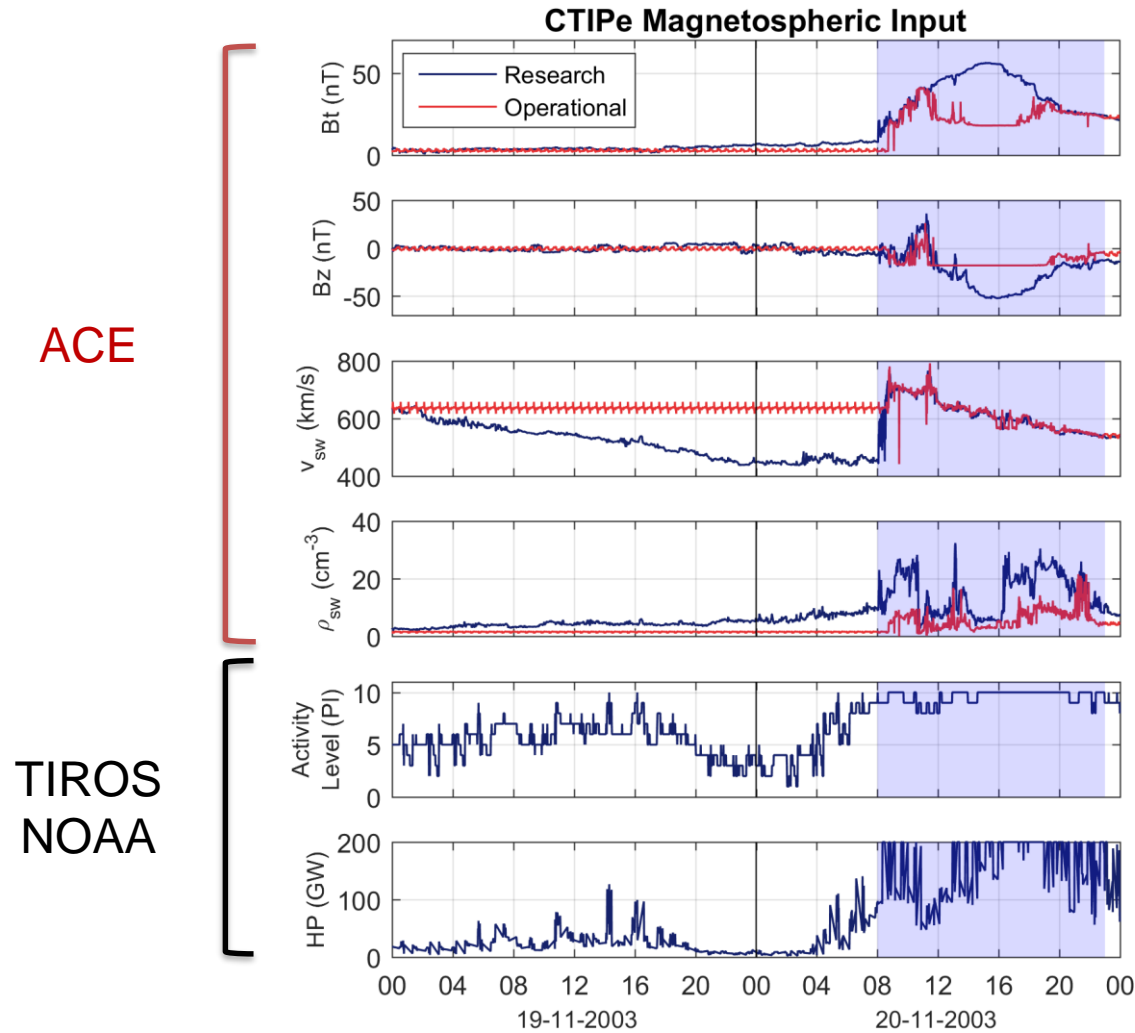


EQ

- **Dst** as indicator of storm phases with an Onset at 8UT and main phase 8-23UT with a minimum -470 nT.
- **AE** shows heating enhancement during the main phase
- **Kp** values increase up to 9 indicating very strong disturbance.



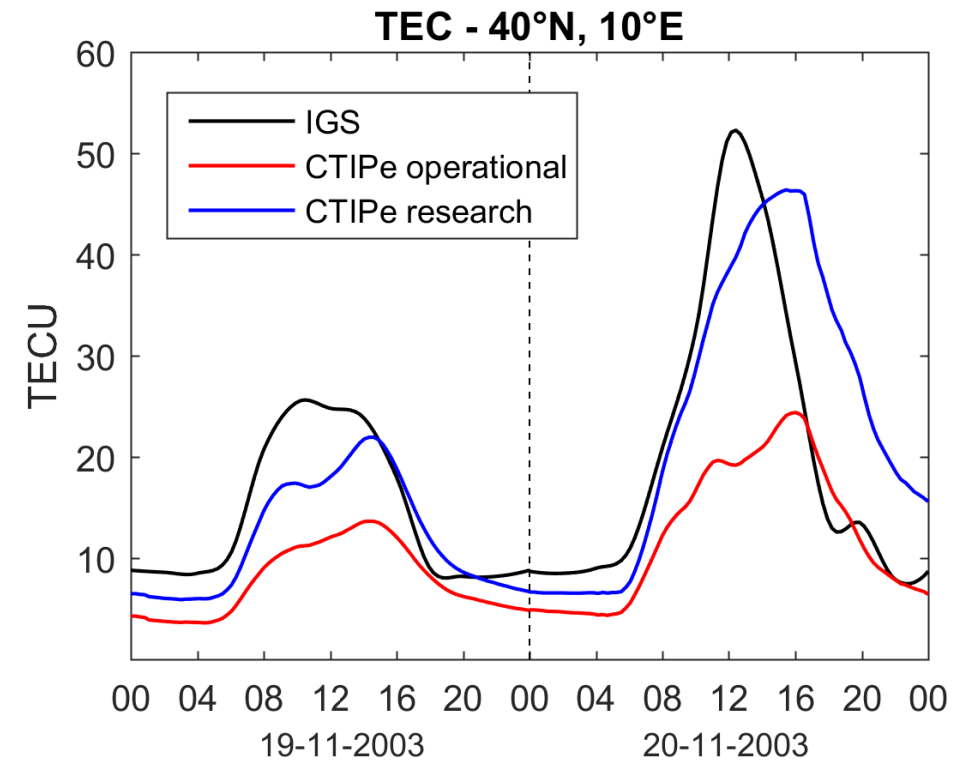
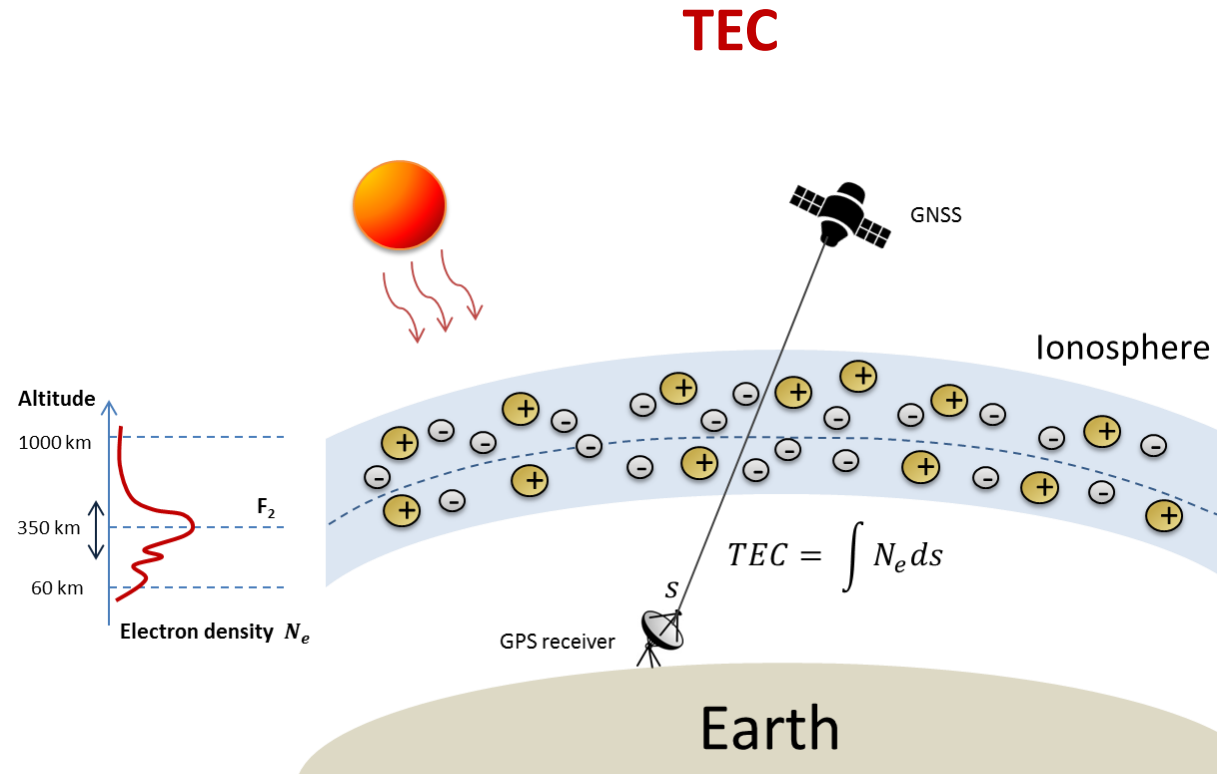
# Operational vs. Research: Input data



- CTIPe magnetospheric input
- **Operational:** Data available in real time
- **Research:** corrected with OMNI data.
- Clear discrepancies **Operational** vs **Research**.

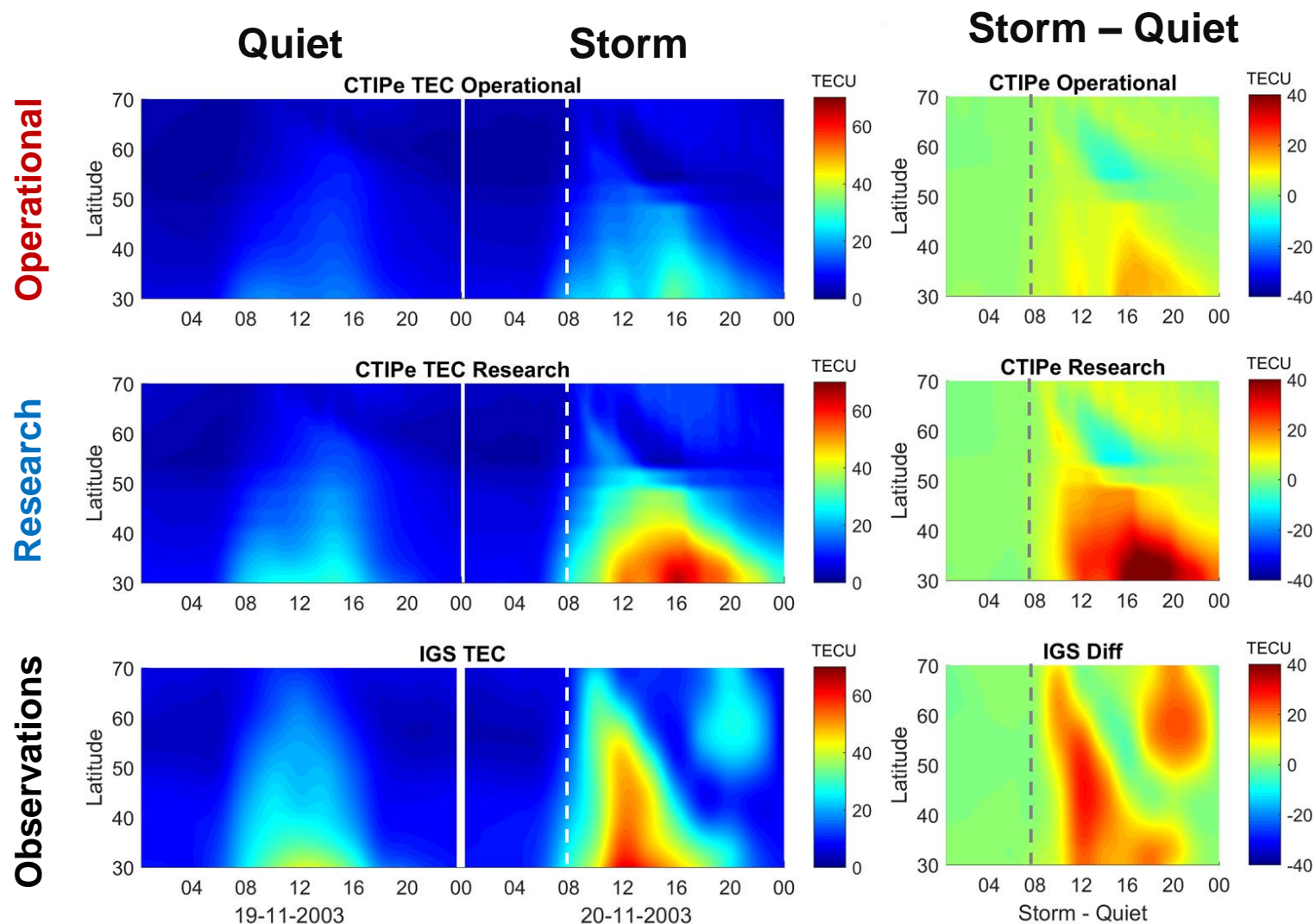


# Operational vs. Research: Ionospheric TEC





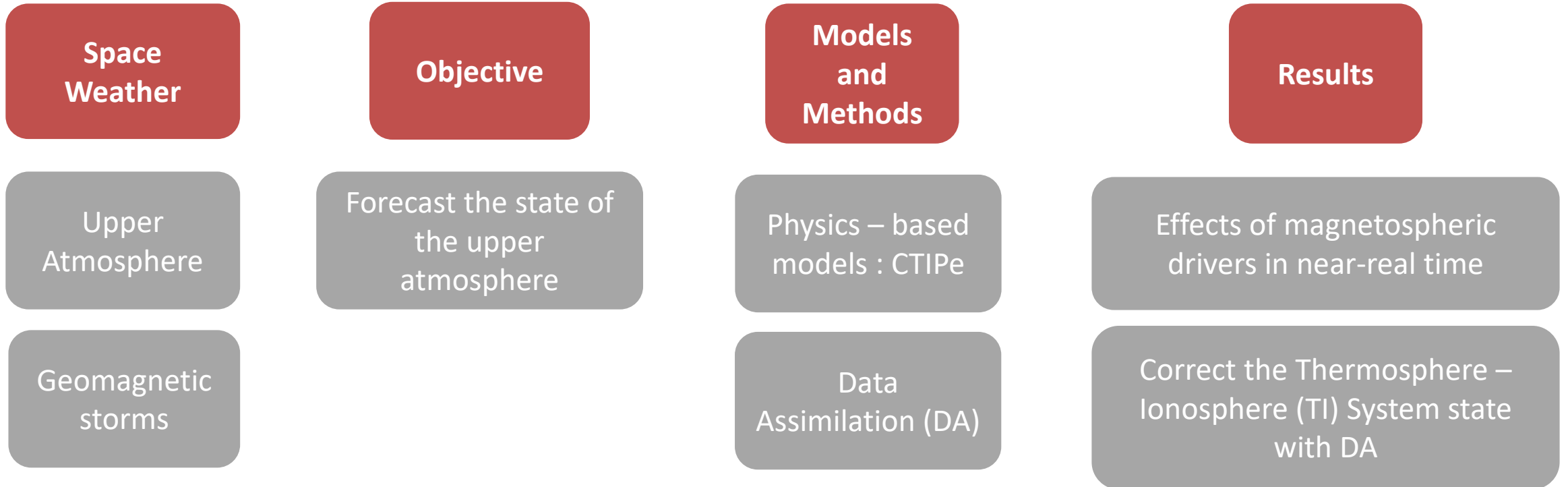
# Operational vs. Research: Ionosphere (TEC) – Europe (10E Ion)



- TEC change in lat vs. time (10E Ion)
- **Operational/Research** runs and IGS TEC observations.
- Onset of the storm (dashed line)
- Enhancement starts high latitudes traveling equatorward with time.
- Discrepancies:
  - **Operational** underestimates and **research** overestimates it.
  - Depletion area travels faster for observations.
  - High latitude evening TEC enhancement.

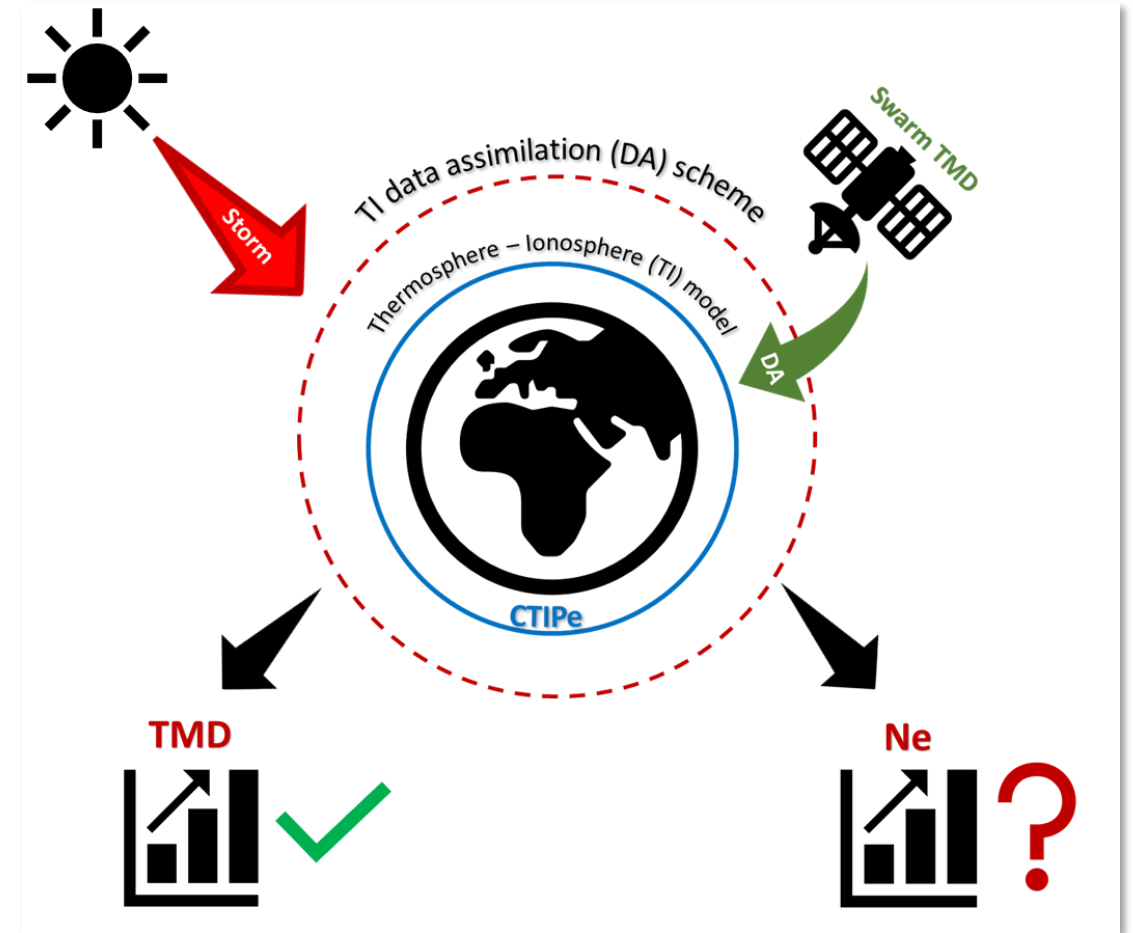


# Outline

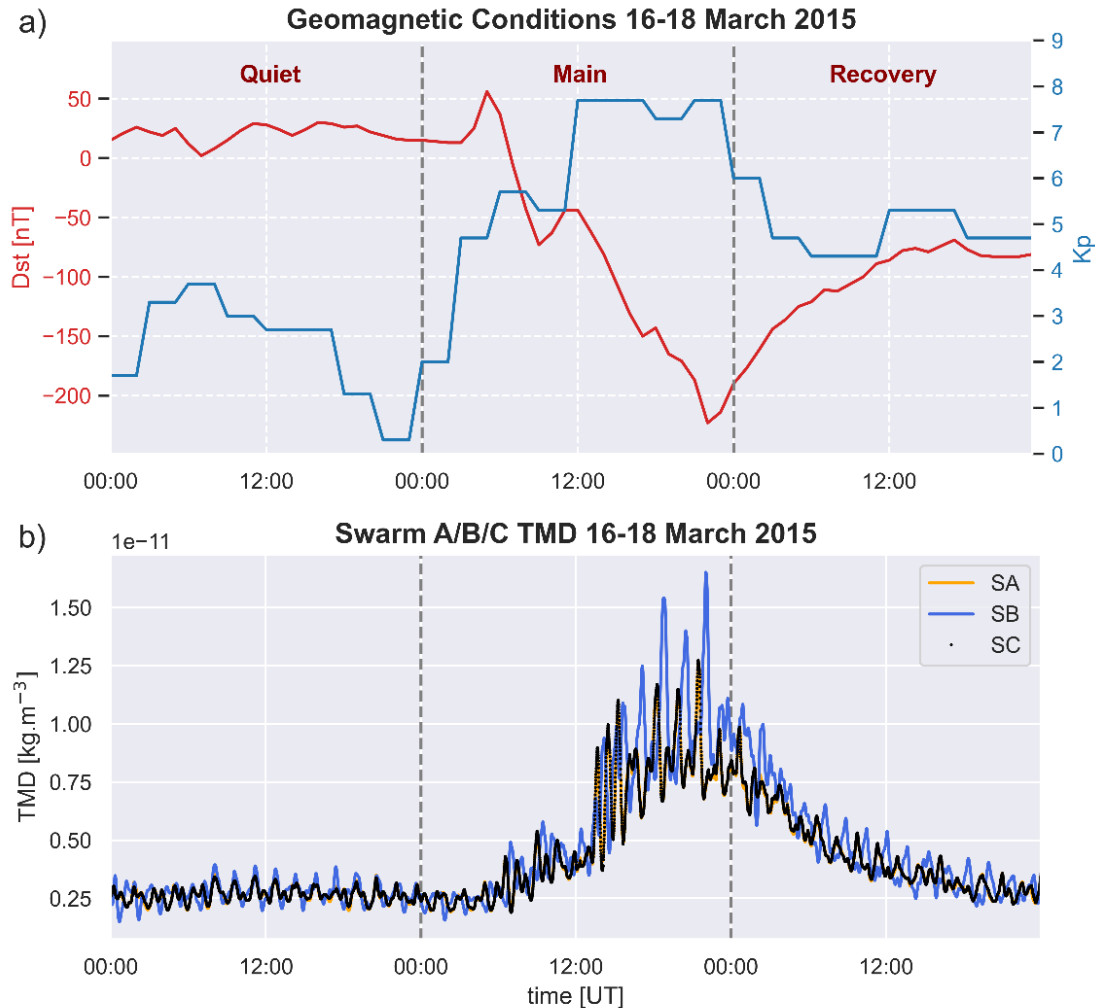


# Neutral density data assimilation effect on electron density

- Can we improve **ionosphere** by assimilating **thermospheric mass density** (TMD) during geomagnetic storm conditions?
  - St. Patrick's Day storm 2015
  - Assimilated Swarm TMD into CTIPe – TIDA
  - Evaluate the Thermosphere – Ionosphere effects



# St. Patrick's Day storm 2015 and DA conditions

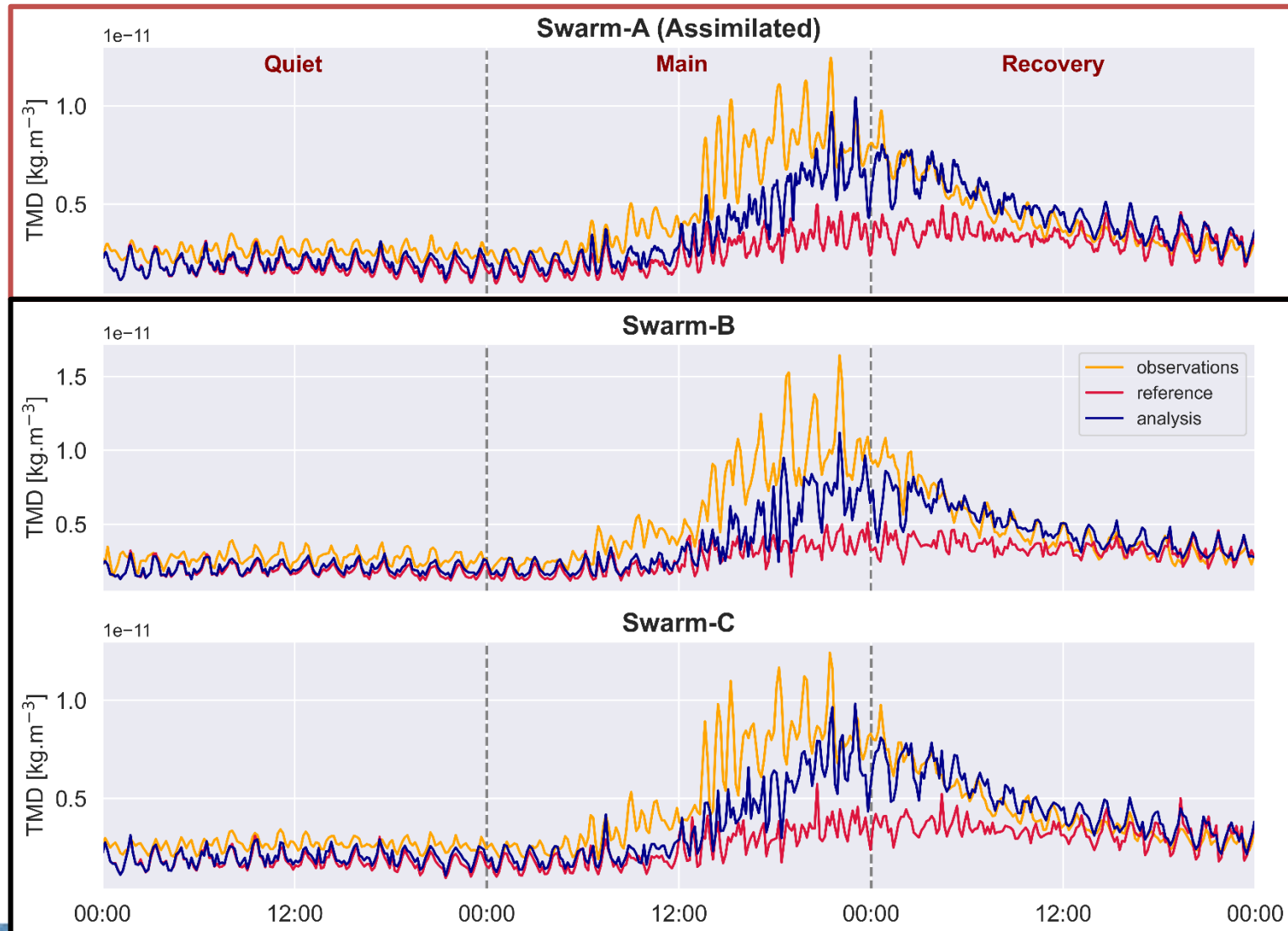


- **Period:** 16-19 March 2015 containing St. Patrick's Day storm
- Days are classified as quiet (16), main phase (17) and recovery (18)
- **TMD Data:** Swarm – A /B/C observations normalized to the common altitude of 400 km.
- **State vector:** Updates the forcing parameters and the necessary quantities to calculate neutral density.
- **Assimilation window:** 10 minutes
- TMD uncertainty is 10%



# Swarm TMD assimilation into CTIPe – TIDA

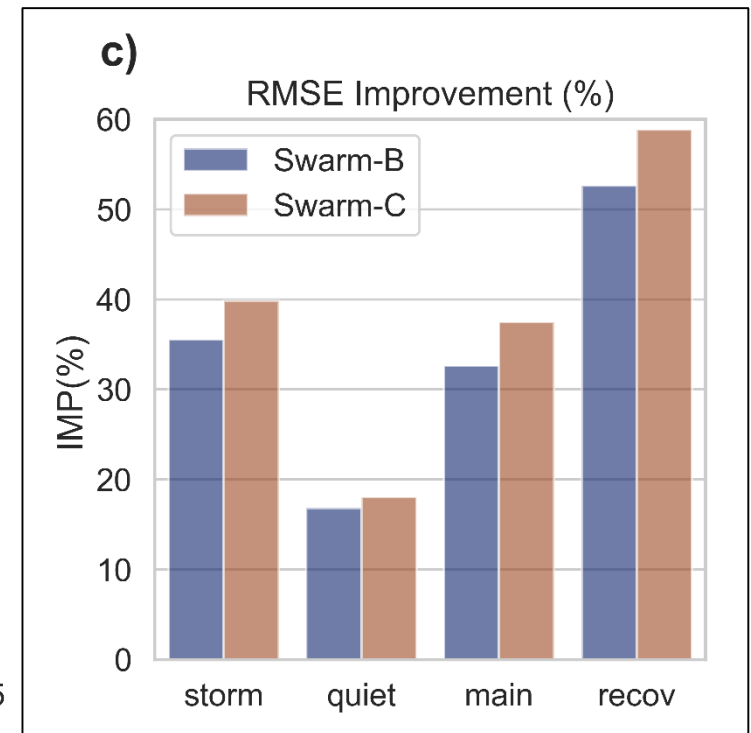
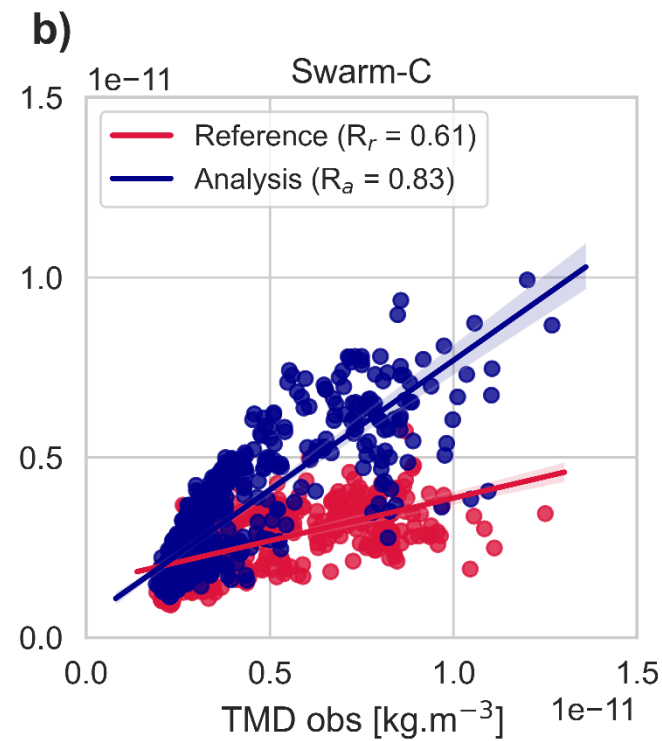
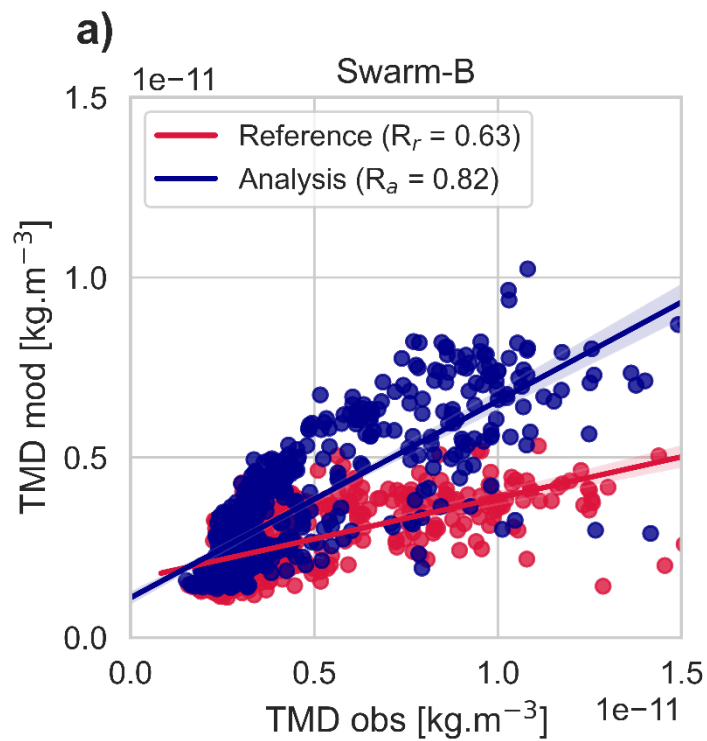
- **Assimilated data:** Swarm – A TMD observations normalized to the common altitude of 400 km.
- **Period:** 16-19 March 2015 containing St. Patrick's Day storm.
- Along the orbit neutral density
- **Observations:** Neutral density from Swarm A/B/C
- **Reference:** Background model results without assimilation
- **Analysis:** Assimilation estimate
- **Differences between Reference and Analysis** show the effect of data assimilation.



# TMD DA impact on the Thermosphere

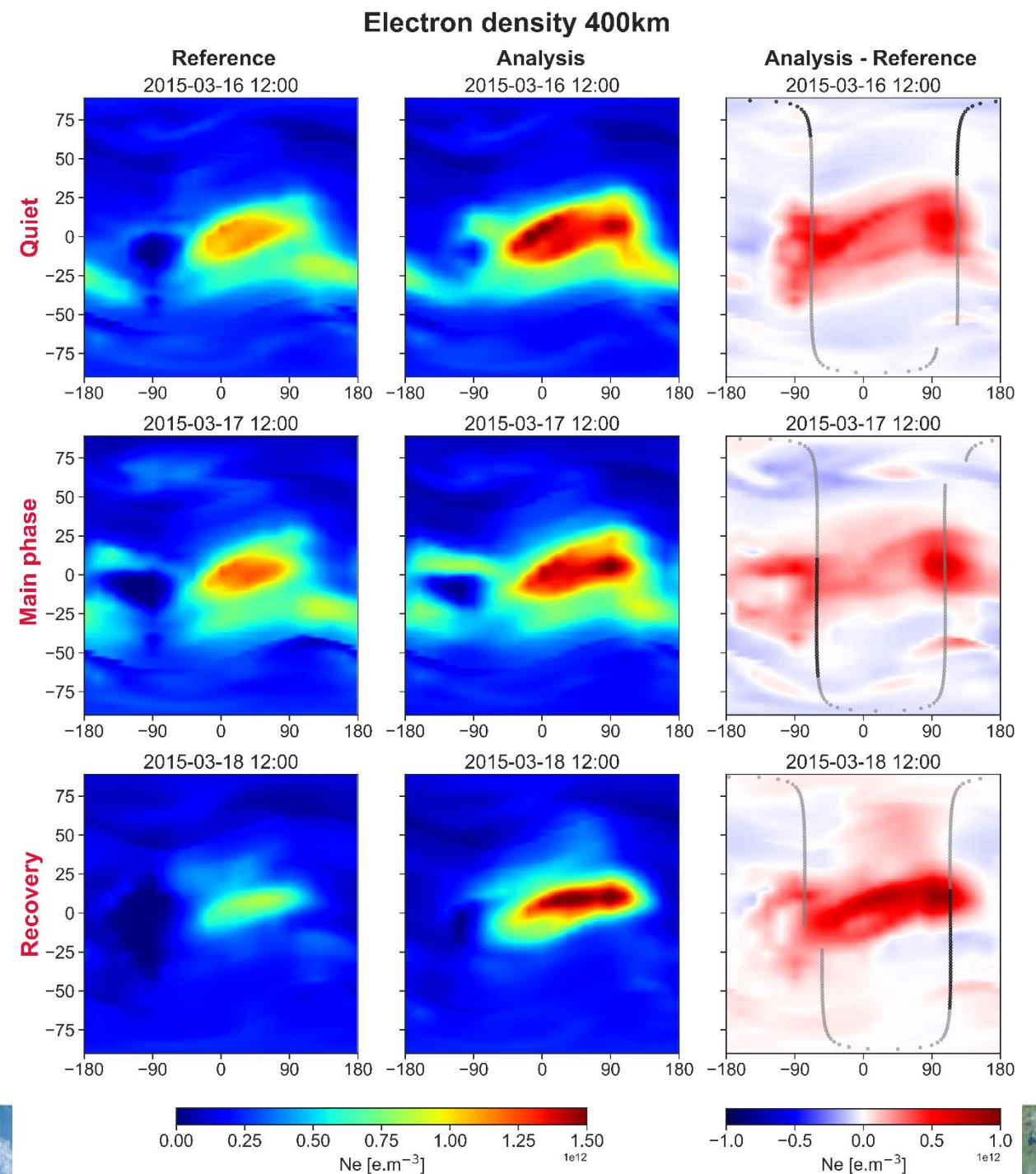
$$\text{RMSE} = \sqrt{\frac{\sum(\text{Obs} - \text{Mod})^2}{N}}$$

$$\text{IMP}(\%) = \frac{(\text{RMSE}_r - \text{RMSE}_a)}{\text{RMSE}_r} 100$$



# TMD DA impact on the Ionosphere

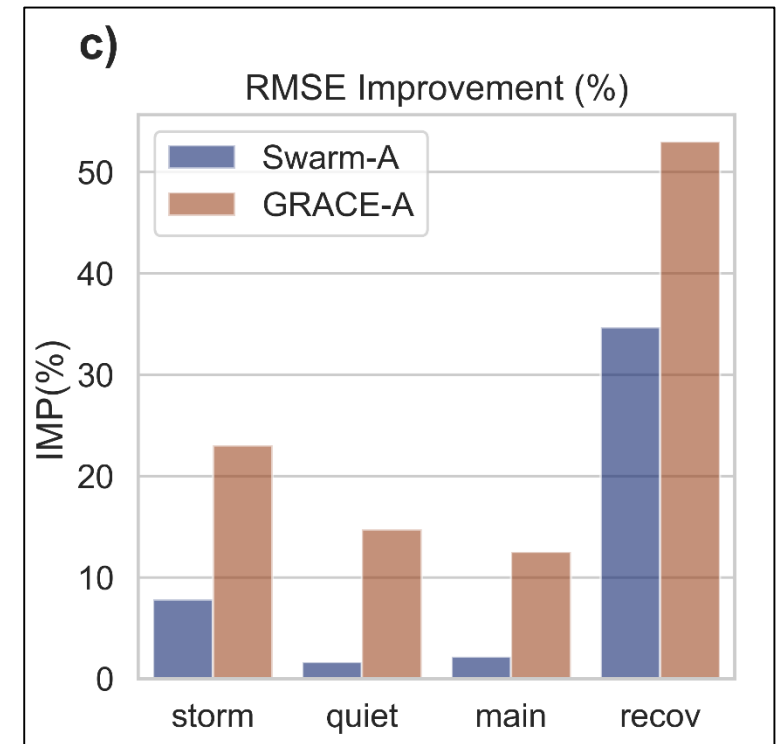
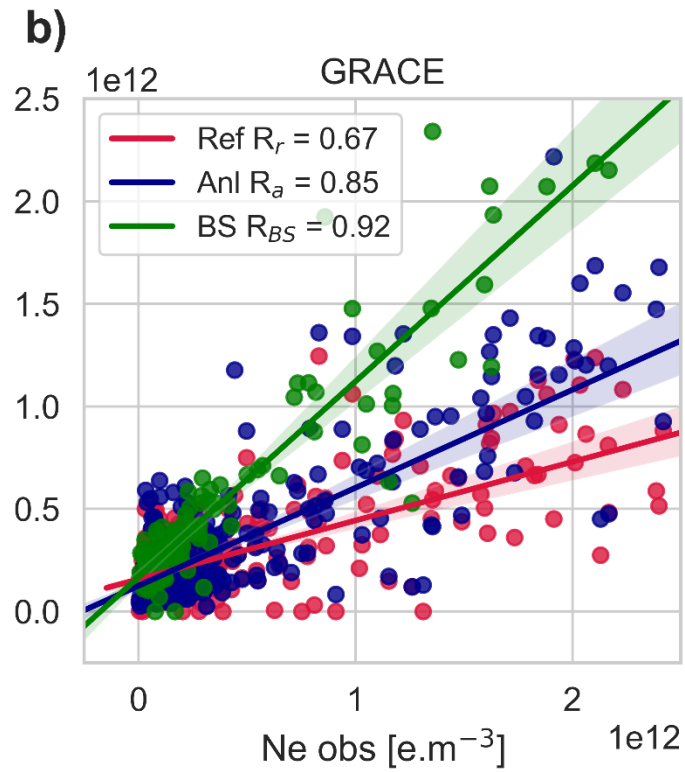
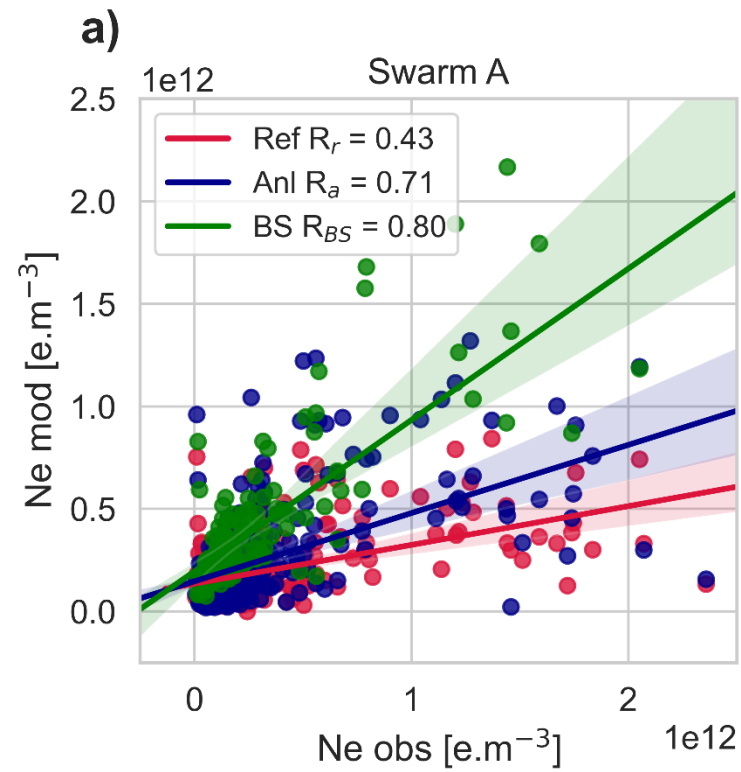
- **Electron density** maps at 400 km altitude
- Reference, analysis and difference
- Quiet day, main and recovery phases at 12:00 UT
- Location of Swarm orbit is represented in the difference plot (grey line)
- Highlighted area of the last two assimilation intervals before 12:00 UT
- The **difference between analysis and reference** shows the **effect of TMD DA in electron density**
- The effect in altitude extends from 200 km to 800 km



# TMD DA impact on the Ionosphere: B-Spline Electron density model

$$\text{RMSE} = \sqrt{\frac{\sum(\text{Obs} - \text{Mod})^2}{N}}$$

$$\text{IMP}(\%) = \frac{(\text{RMSE}_r - \text{RMSE}_a)}{\text{RMSE}_r} 100$$

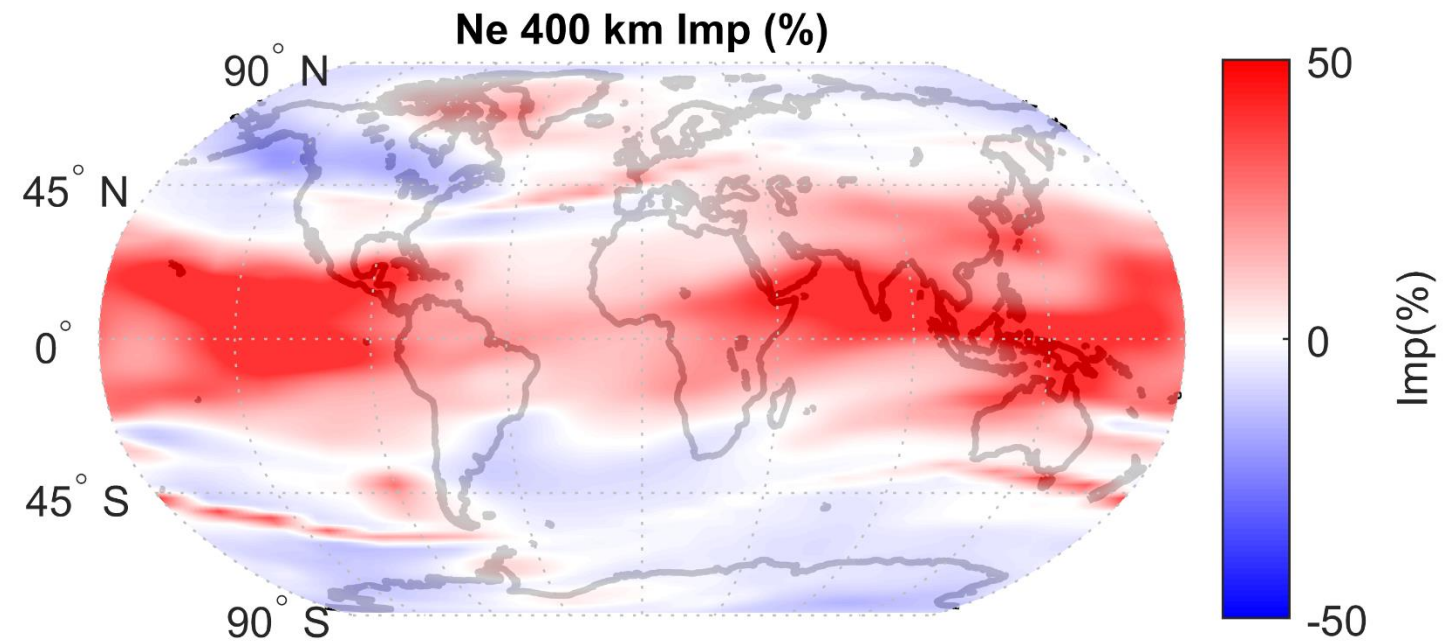




# TMD DA effect on the Ionosphere: B-Spline Electron density model

- **Electron density global improvement at 400 km** between analysis and reference with respect to the B-Spline electron density model
- For the three days of the storm
- Lower RMSE → Better fit of the model to observations
- Improvement (%) of RMSE of the analysis and reference differences.
- **Positive values** are areas of improvement
- The **main area of improvement (red)** is **around the equatorial region (-45, 45) deg latitude**.
- In altitude the improvement extends from 200 km up to 600 km.

$$\text{IMP}(\%) = \frac{(\text{RMSE}_r - \text{RMSE}_a)}{\text{RMSE}_r} 100$$



# Summary

- Discrepancies between model and measurements increase with the increase in **uncertainties of the magnetospheric input data**.
  
- **Assimilation of neutral density** measurements into a physics-based model **during storm conditions** is capable of **correcting the thermosphere and the ionosphere** (with limitations).

*“On the difference between real-time and research simulations with CTIPe”*

I. Fernandez-Gomez, M. Fedrizzi, M. Codrescu, C. Borries, M. Fillion and T. J. Fuller-Rowell.

<https://doi.org/10.1016/j.asr.2019.02.028>

*“Improving estimates of the ionosphere during geomagnetic storm conditions through assimilation of thermospheric mass density”*

I. Fernandez-Gomez, T. Kodikara, C. Borries, E. Forootan, A. Goss, M. Schmidt and M. Codrescu

<https://doi.org/10.21203/rs.3.rs-1342228/v1>



**Thanks for you attention!**

