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CALIBRATION OF A TEMPERATURE-RADIATION EVAPOTRANSPIRATION  
EQUATION FOR UTAH

by

Payam Foroughi

A thesis submitted in partial fulfillment of the  
requirements for the degree

of

MASTER OF SCIENCE

in

Irrigation Science

Approved:

---

Dr. Robert W. Hill

---

Dr. Lyman S. Willardson

---

Mr. Gary P. Merkley

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Dean of Graduate School

UTAH STATE UNIVERSITY  
Logan, Utah

1985



## ACKNOWLEDGEMENT

I wish to express my deepest gratitude towards my parents for their emotional and financial support during my long stay in graduate school. I also thank my advisor Dr. R. W. Hill for his academic assistance. My appreciation also goes to the many new friends from Utah and around the globe that I was fortunate to meet while attending Utah State University.

This work is dedicated to the oppressed people of Iran and others in the world living under dictatorial regimes.

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Two reference crop evapotranspiration (ETr) equations were calibrated to the reference crop of alfalfa using 43 meteorological datafiles from 13 scattered sites in Utah and four sites in the neighboring states of Idaho and Wyoming.

The calibrations were done against the Kimberly version of the Penman ETr formula. The first equation required measured inputs of mean air temperature and solar radiation. It was referred to as the Temperature-Radiation ETr equation ( $ETr = CT \times T \times R_s$ ). The second equation required measured inputs of mean minimum and mean maximum air temperatures. It was referred to as the New Hargreaves equation ( $ETr = K \times T \times TD^{0.5} \times R_s$ ). CT and K

## ABSTRACT

Calibration of a Temperature-Radiation  
Evapotranspiration Equation for Utah

by

Payam Foroughi, Master of Science

Utah State University, 1985

Major Professor: Dr. Robert W. Hill

Department: Agricultural and Irrigation Engineering

Two reference crop evapotranspiration (ETr) equations were calibrated to the reference crop of alfalfa using 45 meteorological datafiles from 13 scattered sites in Utah and four sites in the neighboring states of Idaho and Wyoming.

The calibrations were done against the Kimberly version of the Penman ETr formula. The first equation required measured inputs of mean air temperature and solar radiation. It was referred to as the Temperature-Radiation ETr equation ( $E_{Tr} = CT \times T \times R_s$ ). The second equation required measured inputs of mean minimum and mean maximum air temperatures. It was referred to as the New Hargreaves equation ( $E_{Tr} = K \times T \times TD^{0.5} \times R_a$ ). CT and K



coefficients were found by minimizing a certain objective function which took into account 5-day sums and seasonal sums of ETr. General formulas were derived for estimating CT and K coefficients for any site in the intermountain west area for which the longitude, latitude, elevation, mean longterm July minimum temperature and mean longterm July maximum temperature were known. The average CT and K found for the study sites in Utah were 0.00999 and 0.001074 respectively.

methodologies have been proposed.

(108 pages)

In Utah, methods of estimating agricultural crop water requirements have been introduced to the farmers through the Utah State University Extension Service. The two most recent estimating techniques have been the usage of 1) an Apple micro-computer program model for prediction of 7 and 14 day water use for crops dominant in various areas of Utah, and 2) a miniature, battery operated, field evapotranspiration computer called the "Datapod" (model DF219, Omnidata International, Logan, Utah).

The Datapod is useful for remote sites in Utah and possibly throughout the world for calculation of potential or reference evapotranspiration, where daily weather data are not available. It can store up to 255 days of data without any monitoring by an operator. Figure 1 shows a

## INTRODUCTION

### Need for Study

Expanding world population and constant depletion of world resources forces planners and agriculturalists to find more efficient and productive methods of agriculture and water utilization. As a result crop irrigation requirements have been scientifically scrutinized and various methodologies have been proposed.

In Utah, methods of estimating agricultural crop water requirements have been introduced to the farmers through the Utah State University Extension Service. The two most recent estimating techniques have been the usage of 1) an Apple micro-computer program model for prediction of 7 and 14 day water use for crops dominant in various areas of Utah, and 2) a miniature, battery operated, field evapotranspiration computer called the "Datapod" (model DP219, Omnidata International, Logan, Utah).

The Datapod is useful for remote sites in Utah and possibly throughout the world for calculation of potential or reference evapotranspiration, where daily weather data are not available. It can store up to 255 days of data without any monitoring by an operator. Figure 1 shows a

picture of the Datapod.

The formula used by the Datapod for the calculation of reference evapotranspiration is a version of the Jensen-Haise equation:

$E_{Tr} = CT(T - TX)R_s \times CF$ , where

$E_{Tr}$  is an alfalfa reference crop evapotranspiration (inches/day).  $T$  is the average air temperature (degrees F).  $R_s$  is incoming shortwave solar radiation ( $\text{cal}/\text{cm}^2/\text{day}$ ).  $CF$  is the necessary factor for converting  $E_{Tr}$  from units of  $\text{cal}/\text{cm}^2$  to inches and is equivalent to 0.000673 at 68 °F.  $CT$  and  $TX$  are formula coefficients which can vary from place to place.

The above formula is very simple to use since it only requires the measurement of temperature and solar radiation. However the question can arise as to the reason for calibrating such an equation which may at times not give as accurate a result as more complex methods such as the Penman  $E_{Tr}$  equation.

Usage of a simple method of  $E_{Tr}$  estimation is justifiable under the following arguments:

- Simpler methods are much more feasible for rural areas and developing nations because of: a) less financial burdens on the farmer or the extension



FIG. 1.- Datapod models Similar in Configuration to the DP219 Model which Requires a TP10v Temperature Probe and a Licor Pyranometer. From Ominidata International (27).

#### Objectives

The objectives of this study are :

of affairs at selected sites in the intermountain west area. This will be done by finding appropriate CT and TX coefficients for various sites throughout Utah and some sites in the neighboring states of Idaho, and Wyoming.

agency, b) possible lack of technical and scientific know-how which more complex methods require, and c) ease in availability of simple meteorological data. In short, a simple method of ETr estimation provides a methodology which is "user" oriented rather than "research" oriented (22).

3) To test a method of ETr estimation for  
— The estimation of ETr is recommended for decision making in irrigation scheduling. The risk of missing data increases with complexity of method and instrumentation, and can lead to incorrect decision making in water management.

### Objectives

The objectives of this study are :

1) To calibrate a temperature-radiation ETr formula to the reference crop evapotranspiration of Alfalfa at selected sites in the intermountain west area. This will be done by finding appropriate CT and TX coefficients for various sites throughout Utah and some sites in the neighboring states of Idaho, and Wyoming.

2) To analyse the meteorological and site factors among the study sites which can cause distinct pairs of CT and TX coefficients for each site. This can tentatively lead to a generalized procedure for finding appropriate coefficients for any site in Utah and surroundings.

3) To test a method of ETr estimation for agricultural locations in Utah whose only available meteorological parameters are maximum and minimum temperatures. The method to be tested is the New Hargreaves ET equation developed at the Utah State University International Irrigation Center.



## LITERATURE REVIEW

A general estimation is that 80% of water consumed worldwide is attributable to irrigation (18). The increasing demand for irrigation water has caused engineers and planners to come up with ingenious methods of water utilization. In California, giant canals have been constructed for water delivery to the fertile soils of the San Joaquin valley. Soviet planners are considering the construction of a huge network of dams and canals to reverse the flow of the Pechora and Ob rivers that now rush uselessly into the Arctic sea. The potential waters are to be made available for irrigation in southern Russia (11). In a similar effort, the Chinese government is building canals to divert the Chang Jiang river 750 miles to the North China Plain (36). Cost has been estimated at \$13.2 billion. Scientists at the University of California, Davis, have developed strains of wheat, barley and tomatoes which can be irrigated by up to 70% pure seawater. The strains may be capable of creating productive agricultural regions out of millions of acres of the world's sandy deserts (34).

Water consumption by man and by nature through a variety of means are identified by specific terminologies. The combination of the evaporation of water from soil and crop surfaces and the transpiration through crop stomates is termed "evapotranspiration". Penman (28) initiated the term "ET", but for it not to mean evapotranspiration, rather to indicate "evaporation from turf". He writes:

Many find it helpful to give a special name to the combined effect, referring to it as 'evapotranspiration'. ...it is presumably a useful term, but it is rather ugly[!], and it hardly seems necessary, as there are few situations in which the use of 'evaporation' or 'transpiration' is not entirely adequate (37, p.54).

Contemporary irrigation engineers and agriculturalists, however, refer to "ET" as an abbreviation for evapotranspiration. Many use the term "potential evapotranspiration",  $ET_p$ , or "reference crop evapotranspiration",  $ET_r$ . The term  $ET_p$  is usually used in conotation with the amount of water evapotranspired in a unit of time by a green crop, completely shading the ground, of unifrom height and never short of water (28).

The actual bio-physical process of ET cools the crops and protects them from overheating. This is done by the transfer of 85 calories to the surrounding air for every gram of water evaporated. Dew formation at night and early morning is the reverse process in that it acts as



insulation for crop surfaces by making available to the crop the energy resulting from water vapor condensation (37). The ET process also helps in conducting the nutrient solution necessary for crop growth and development to various parts of the crop. Studies show that ET is directly related to crop yields. The majority of yield-ET studies have found a linear relationship between the two (10,15); for example, Hill (15) found the following yield-ET relationship for southern Utah:  $Y = 0.243(ET) - 0.765$  where Y is alfalfa yield in tons/acre, and ET is in inches.

ET and crop growth are known to be regulated by plant, soil, and climatic factors. With a relatively constant effect by the plant and soil factors, one should be able to estimate ET through climatic data. Historically scientists have not been easily convinced of the close relation among ET, yield, and meteorological parameters. In 1871 Koppen referred to the relation between climatic factors and crop development as:

...the beautiful illusion that it is possible to represent the development of plants, even those of a single species, by means of a general formula which contains temperature, light, humidity and other external agents as factors. But, to be sure, whoever finds illusion more pleasant than sober knowledge is not disturbed by such considerations; so he goes on his way in peace and it is no fault of his if others can not follow him (38, p.459).

Abbe wrote in 1905:

Ofcourse, hydraulic and irrigation engineers need to know the loss of water by evaporation, but in nature this is so mixed up with seepage, leakage, and consumption by animals and plants that our meteorological data are of comparatively little importance (1, p.254).

With more scientific scrutiny the relationship of ET with meteorological factors became more clear. In 1915 at Akron, Colorado, Briggs (5) conducted a thorough study to determine the relationship of ET from a variety of crops to different weather factors. For most crops he found a high degree of correlation between ET and wet-bulb temperature followed by air temperature and solar radiation. Since then a plethora of methods, models, and equations for the quantitative estimation of evapotranspiration have been proposed. Usage of any specific method will depend on the type of meteorological data available and the accuracy of that method in estimating ET. The following describes some ET estimation methods and the research done by scientists in testing those methods. Emphasis is put on the temperature-radiation methods and the Penman combination equation.

#### Temperature-Radiation Methods

In 1961, the French scientist Turc (39) proposed an equation for estimation of ET<sub>p</sub> with the only inputs of

mean air temperature and solar radiation. He worked with data from France, Denmark, Ireland, Morocco, Tunisia, Congo, Iraq, and Ceylon to derive the following formula:

$$ET_p = 0.013 [T / (T + 15)] (R_s + 50) \dots \dots \dots (1)$$

where  $ET_p$  is in mm/day,  $T$  is mean air temperature in degrees C, and  $R_s$  is the incident solar radiation in equivalent units of mm/day. The above equation is claimed by its author to give satisfactory results for areas with mean relative humidities (RH) of 50% or higher. For more arid regions with RH values less than 50%, Turc offers the following:

$$ET_p = 0.013 [T / (T + 15)] (R_s + 50) [1 + (50 - RH) / 70] \dots \dots \dots (2)$$

In 1963 Jensen and Haise (19) studied the evapotranspiration process based on the energy balance concept. The following relationship was proposed:

$$R_s(1-r) - Ret - ET - G - A \approx 0 \dots \dots \dots (3)$$

where,  $R_s$  = shortwave solar radiation flux,  
 $r$  = crop reflectance or albedo,  
 $Ret$  = effective or net thermal longwave radiation,  
 $ET$  = rate of evapotranspiration,  
 $G$  = sensible heat flux to or from the ground, and  
 $A$  = sensible heat flux to or from the air.

Equation (3) was expressed in dimensionless form when divided by  $R_s$ :

$$ET/R_s = 1 - r - Ret/R_s - G/R_s - A/R_s \dots \dots \dots (4)$$

The ratio  $ET/R_s$  was said to represent the combined effects of net thermal radiation, heat flux to or from soil and air, and other minor components. Because of the non-linear relationship between saturated vapor pressure and air temperature and its effect on thermal radiation, air temperature was to have a direct effect on the  $ET/R_s$  ratio provided adequately irrigated crops, where evaporating and transpiring surfaces do not limit the vaporization of water. Working with data of approximately 1000 individual sampling periods during the growing season of fifteen crops obtained within a span of 35 years from four different climatic regions in the U.S., Jensen and Haise found a general linear relationship with good regression coefficient between the  $ET/R_s$  ratio and mean air temperature. Figure 2 shows the regression line and the data points used (19). The relationship found came to be known as the original Jensen-Haise evapotranspiration equation and is as below:

$$ET_p = 0.014(T - 26.4)R_s \dots \dots \dots (5)$$

where,  $ET_p$  = potential evapotranspiration (in/day),

$T$  = mean air temperature (deg. °F), and

$R_s$  = incident solar radiation (equivalent of in/day).



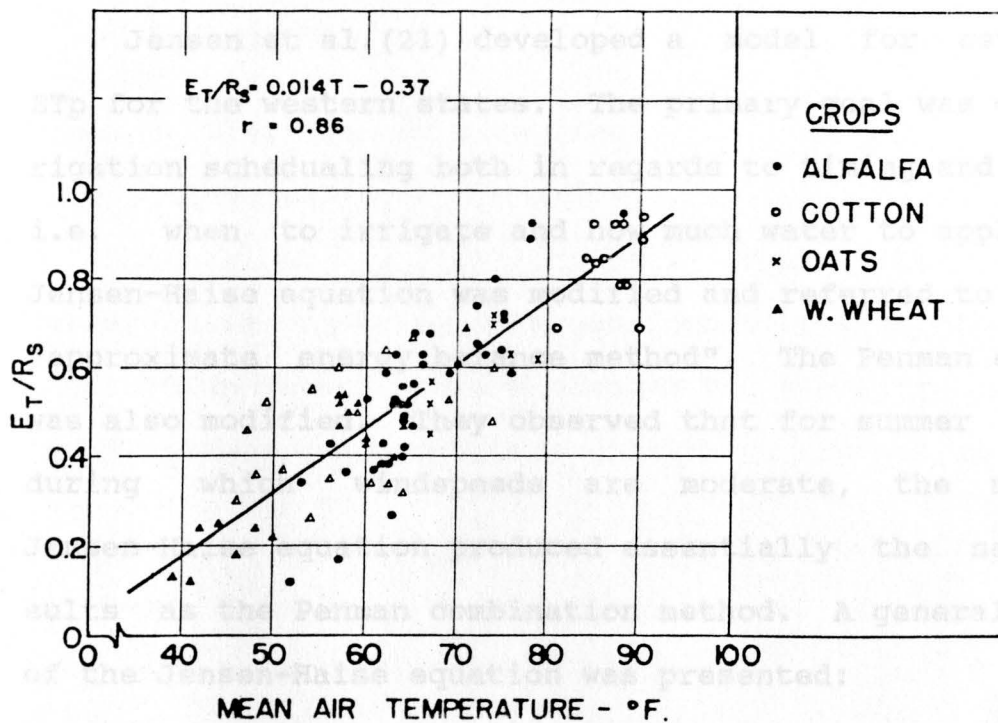


FIG. 2.- The Original Jensen-Haise Linear Regression Between  $E_T/R_s$  and Mean Air Temperature,  $T$ . From Jensen and Haise (19).

$$E_T = C_T(T - T_x)R_s \dots \dots \dots (6a)$$

$$C_T = 1 / (C_1 + 13C_h) \dots \dots \dots (T \text{ in deg. } ^\circ F) \dots \dots \dots (6b)$$

$$C_T = 1 / (C_1 + 7.3C_h) \dots \dots \dots (T \text{ in deg. } ^\circ C) \dots \dots \dots (6c)$$

where  $C_h$  is a humidity index and  $C_1$  is an elevation index.

$$C_h = [37.5 \text{ mmHg} / (e_1 - e_2)] = [50 \text{ mb} / (e_1 - e_2)] \dots \dots \dots (6d)$$

Equation (5) was claimed by its authors to estimate good results of ETP for arid and semiarid areas.

Jensen et al. (21) developed a model for estimating ETP for the western states. The primary goal was good irrigation scheduling both in regards to timing and amount, i.e. when to irrigate and how much water to apply. The Jensen-Haise equation was modified and referred to as the "approximate energy balance method". The Penman equation was also modified. They observed that for summer months, during which windspeeds are moderate, the modified Jensen-Haise equation produced essentially the same results as the Penman combination method. A general format of the Jensen-Haise equation was presented:

$$ETp = CT(T - TX)Rs \dots \dots \dots (6a)$$

The CT and TX coefficients were proposed to be derived for specific sites as so:

$$CT = 1 / (C_1 + 13C_h) \dots \dots \dots (T \text{ in deg. } ^\circ F) \dots \dots \dots (6b)$$

$$CT = 1 / (C_1 + 7.3C_h) \dots \dots \dots (T \text{ in deg. } ^\circ C) \dots \dots \dots (6c)$$

where  $C_h$  is a humidity index and  $C_1$  is an elevation index.

$$C_h = [37.5 \text{ mmHg} / (e_1 - e_2)] = [50 \text{ mb} / (e_1 - e_2)] \dots \dots \dots (6d)$$

$e_1$  and  $e_2$  are saturation vapor pressures (mmHg or mb) during the warmest month of the year at the mean maximum and minimum air temperatures respectively.

$$C_1 = 68 - 3.6E/1000 \dots \dots \dots (6e)$$

where E = elevation of site (feet),

$$TX = 27.5 - 0.25(e_1 - e_2) - E/1000 \dots \dots \dots (6d)$$

Jensen et al. (22) indicate that the modified Jensen-Haise equation is a good substitute for the Penman combination method where windspeed and humidity data are not available, and advective conditions are not severe.

In 1965 Stephens (35) worked with ET data from Florida, North Carolina and Davis, California. He came up with four regression equations of ET/Rs vs. T for different climatic regions depending on their average July local noon relative humidity. Figure 3 shows a U.S. map as it was divided into four different areas based on mean RH values corresponding to Stephens' classification. Stephens took the original Jensen-Haise coefficients to represent the western and mid-western U.S. with average July noon RH of less than 40%. Table 1 shows his results.

Grabow (9) did regression analysis of measured ET over solar radiation vs. the average temperature for different sites. He found large negative X-intercepts which

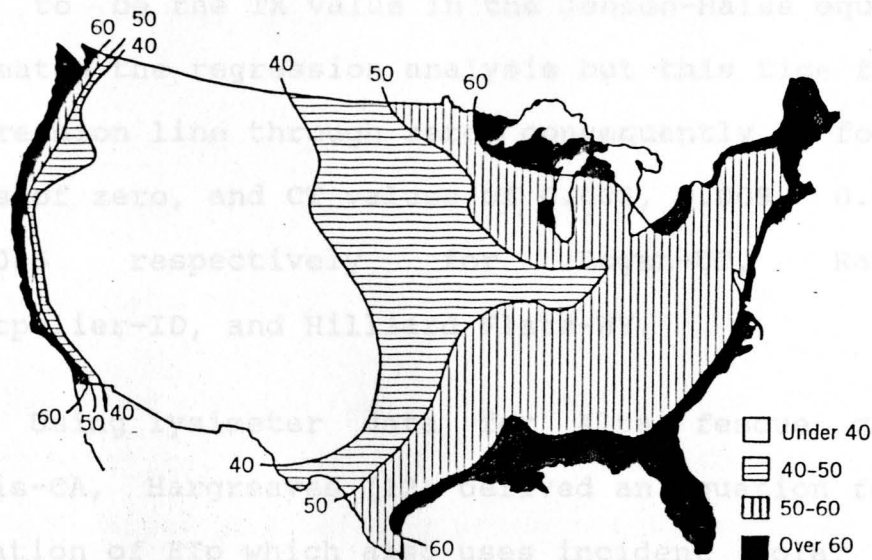


FIG. 3.- U.S. Map as Divided by Stephens into Four Areas of Similar Mean Relative Humidity Values. From Stephens (35).

Table 1.- Stephens' Proposed ET Equations for the Continental U.S. Based on Relative Humidity (RH) Values. From Stephens (35).

Average RH	Regression Equation	CT	TX
RH > 60%	$ET/R_s = 0.00820(T) - 0.1900$	0.00820	23.2
50% < RH < 60%	$ET/R_s = 0.00868(T) - 0.1938$	0.00878	22.1
40% < RH < 50%	$ET/R_s = 0.01067(T) - 0.2256$	0.01067	21.1
RH < 40%	$ET/R_s = 0.01400(T) - 0.3700$	0.01400	26.4



are to be the TX value in the Jensen-Haise equation. He repeated the regression analysis but this time forcing the regression line through zero; consequently he found TX values of zero, and CT values of 0.008, 0.009, 0.0085, and 0.0085 respectively for Logan-UT, Randolph-UT, Montpelier-ID, and Hilliard Flats-WY.

Using lysimeter data for Alta fescue grass from Davis-CA, Hargreaves (12) derived an equation for the estimation of ET<sub>p</sub> which also uses incident solar radiation and mean temperature:

$$ET_p = 0.0075(R_s)T \dots\dots\dots(7)$$

where, ET<sub>p</sub> = potential grass ET (in/day),

Rs = incident solar radiation in equivalent depth of water evaporation,

T = average daily temperature (deg. °F).

Assuming an alfalfa-grass ET<sub>p</sub> conversion factor of 1.2, the 0.0075 in the Hargreaves equation would change to 0.0090; notice that this would be a version of the Jensen-Haise equation for a reference crop of alfalfa with CT of 0.009 and TX of zero.

Hargreaves and Samani (13) estimated the incident solar energy, R<sub>s</sub>, as a function of extraterrestrial radiation, R<sub>a</sub>, and difference of maximum to minimum temperature and daytime wind conditions,

tures, TD. The resulting equation was as follows:

$$R_s = K \times R_a \times TD^{0.50} \dots\dots\dots(8)$$

K is a calibration coefficient. Equation (8) was combined with equation (7), resulting in the relationship below:

$$ET_o = K (R_a \times T \times TD^{0.50}) \dots\dots\dots(9)$$

If the above equation be valid, and if the coefficient K is locally calibrated, it would have very practical usage in rural areas and farming communities, since the only measurements required are maximum and minimum temperatures.

Doorenbos and Pruitt (8) present an ET estimation method in the FAO Irrigation and Drainage Paper #24. It looks as below:

$$ET_o = c(W.R_s) \dots\dots\dots(10)$$

where  $ET_o$  is the reference crop ET in mm/day. It was defined as:

...the rate of evapotranspiration from an extensive surface of 8 to 15 cm tall, green grass cover of uniform height, actively growing, completely shading the ground and not short of water (8, p.9).

$R_s$  is incident solar radiation in equivalent of water evaporation.

W is a weighing factor depending on the temperature and altitude.

c is an adjustment factor depending on the mean humidity and daytime wind conditions.

Doorenbos and Pruitt present tables and graphs to determine specific W and c values which depend on the altitude, mean temperature, mean RH, and mean wind movement of an area.

### Penman Combination Equation

Aristotle is known to have contempolated whether the sun or the wind is the most important factor influencing evaporation. He is known to have chosen wind for he said it carried the vapor away (28). In 1948 Penman (29) indicated that in order for evaporation to happen, there are two necessary requirements. One is the supply of energy better known as the latent heat of vaporization, and the other is a mechanism for removing the water vapor or a sink for vapor. These two processes are known as the "energy balance" and the "turbulent transfer" concepts, and form the bases of the Penman combination equation. Penman's equation in 1948 had the following format:

$$E_o = (Rn \Delta + 0.27E_a) / (\Delta + 0.27) \dots \dots \dots (11)$$

where,  $E_o$  = evaporation from open water (mm/day),

$Rn$  = net radiant energy available at the surface in evaporation equivalent (mm/day),

$\Delta$  = slope of the saturation vapor pressure and temperature ( $^{\circ}F$ ) curve ( $de_a/dT$ ),

$$E_a = 0.35(1+0.0098U_2)(e_a - e_d),$$

$U_2$  = mean wind velocity at 2 meter height (miles/day),

$e_a$  = saturation vapor pressure at the air temperature (mmHg), and

$e_d$  = saturation vapor pressure at the dew-point temperature (mmHg).

Penman observed that crop ET averaged about  $0.75E_o$ .

Later Penman (30,31) modified his equation of open water evaporation to represent evapotranspiration as such:

$$ET = [(\Delta / \gamma)R_n + E_a] / [(\Delta / \gamma) + 1] \dots \dots \dots (12)$$

where, ET = potential crop evapotranspiration (mm/day),

$\gamma$  = wet and dry-bulb psychrometric equation constant,

$E_a$  is an expression for the "drying power" of the air involving windspeed and saturation deficit:

$$E_a = 0.35(1 + U_2/100)(e_a - e_d), \text{ and}$$

$\Delta$ ,  $u_2$ ,  $e_a$ , and  $e_d$  are as defined earlier.

Penman (28,31) indicated that since the main source of energy for ET is radiant sunshine, the heat or energy reaching the vegetation would be:

$$H = R_s(1-r) - R_b \dots \dots \dots (13)$$

where, H=heat budget or net radiation,  $R_n$ ,

$R_s$ =shortwave incoming radiation,

$r$ =reflection coefficient or albedo,

$R_b$ =net longwave outgoing or back radiation.

He also states that the expenditures for the heat budget are mainly dual. One being ET, and the other being sensible heat transfer to the air A, i.e.:  $H = ET + A$ . If the condition of adequate water supply to vegetation is met, the ratio of A/ET can be a small and constant value. Since the effect of sensible heat transfer, A, is to raise the surrounding air temperature, air temperature measurements can serve as good indicators of A and ET. Under conditions of adequate water supply (or water non-limiting), higher air temperatures can be directly associated with higher ET. However if the condition of adequate water supply is not met, the reverse can be true, i.e. more energy is used for sensible heat, and less for ET. This can cause higher crop canopy and soil temperatures, and lower ET, hence creating an indirect relationship between air temperature and ET. In such conditions air temperature measurement used as a parameter in ET formulas can be misleading, and will over-estimate ET.

Jensen et al.(21) using lysimeters in Kimberly, Idaho, modified and calibrated the Penman equation. The current version of the modified Penman equation which is used extensively throughout the intermountain U.S. stands as such (16):

equation, refer to Appendix I.



$$ETr = [ \Delta / ( \Delta + \gamma ) ] (Rn+G) + [ \gamma / ( \Delta + \gamma ) ] 15.36 (W_1 + W_2 U_2) (e_s - e_a) \dots \dots \dots (14)$$

where, ETr = reference crop ET for "...well watered actively growing alfalfa with sufficient growth for near maximum ET in arid, irrigated regions" (40). ETr is in units of langleys/day or in/day if multiplied by  $1/(585 \times 2.54) = 0.000673$ ,

$\Delta$  and  $\gamma$  are as defined previously,

Rn = net radiation (langleys/day),

G = soil heat flux (langleys/day),

$W_1$  and  $W_2$  are empirical wind parameters dependent on location and type of crop grown,

$U_2$  = wind movement at 2 meters height (miles/day),

$e_s$  = saturation vapor pressure at mean air temperature (mmHg) as in the mean of saturation vapor pressures at mean daily maximum and minimum air temperatures,

$e_a$  = mean actual vapor pressure (mmHg) to be equivalent to the saturation vapor pressure at mean daily dewpoint temperature, Td. Td can be approximated from a single morning dewpoint temperature determination.

The above inputs required for the modified Penman ETr equation can be empirically determined leaving only the following required meteorological measurements:

1. maximum daily air temperature Tmx,
2. minimum daily air temperature Tmn,
3. early morning dew-point temperature Td,
4. incoming solar radiation Rs,
5. daily wind travel  $U_z$ , measured at height z.

For a description and format of the empirical formulas used in the modified Penman combination ET estimation equation, refer to Appendix I.

## FLOW CHART

Estimation of  $E_{Tr}$  is one of the steps necessary for finding the amount of water lost by field crops or the actual crop ET. Usually a crop factor,  $K_c$ , is multiplied by  $E_{Tr}$  to arrive at the actual ET. Figure 4 from Burman et al. (6) demonstrates a systematic approach of estimating ET and irrigation water requirements through different methods.

Wright (40) developed improved crop coefficients for various irrigated crops at southern Idaho. Initially, measured daily alfalfa ET, using lysimeters, were to be used as reference ET in Wright's research for calculation of crop coefficients, while estimated  $E_{Tr}$  from the modified Penman combination method were to be used only when the alfalfa was not at full cover. However, for the eight years of lysimeter data at Kimberly-ID, measured ET was not at a maximum level during much of the growing season. The reasons cited were the time required for the alfalfa to reach full cover in the spring and after each cutting, and the lodging caused by wind and rain. Hence Wright used the modified Penman combination method to find a procedure for calculating the final ET. The Penman method provided a continuous and consistent data base for obtaining  $k_c$  for the entire growing season.

## FLOW CHART

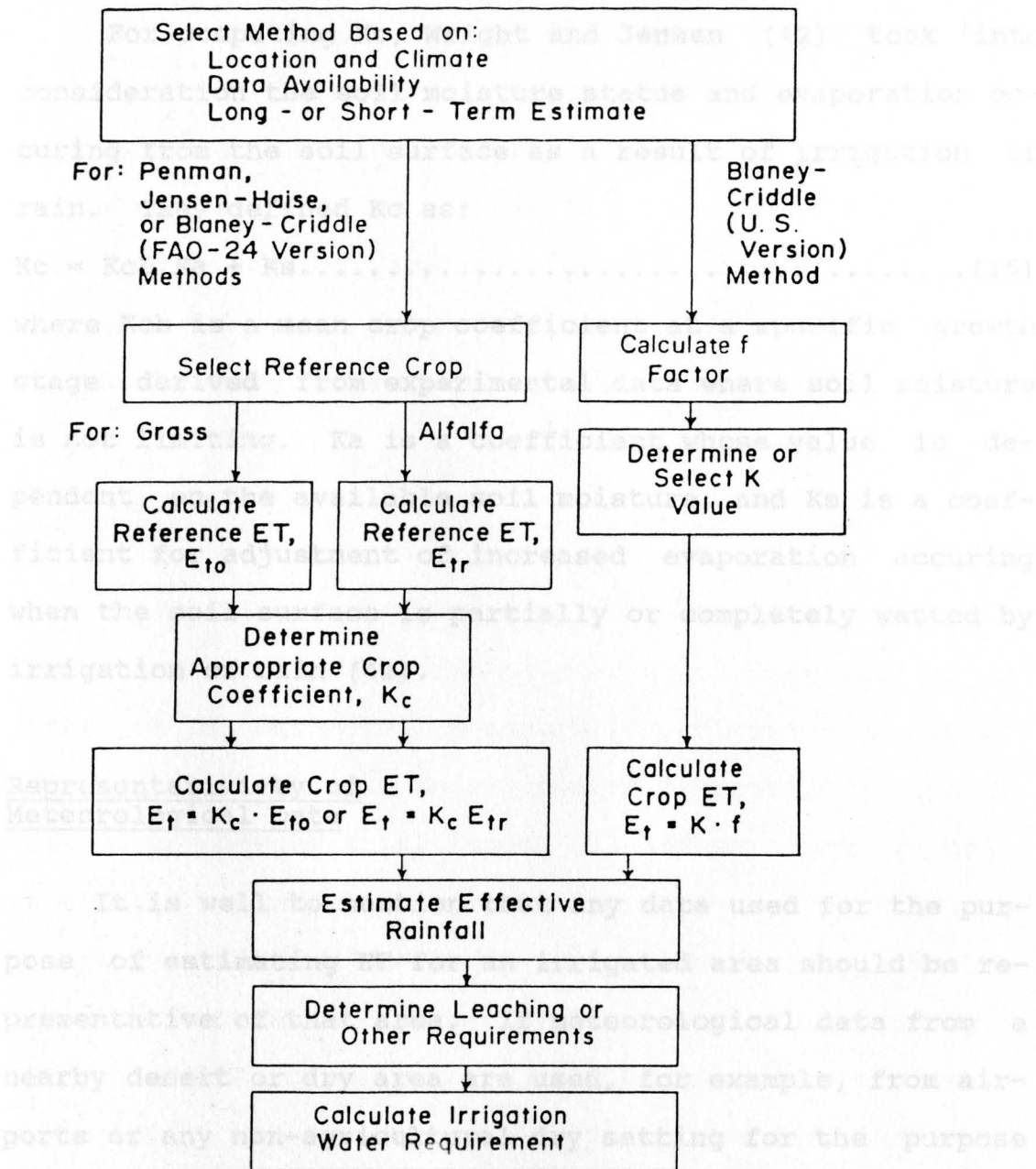


FIG. 4.- Flow Chart Demonstrating the Steps Necessary for the Calculation of Crop ET and Irrigation Water Requirements. From Burman et al. (6).



For computing  $K_c$ , Wright and Jensen (42) took into consideration the soil moisture status and evaporation occurring from the soil surface as a result of irrigation or rain. They defined  $K_c$  as:

$$K_c = K_{cb} \cdot K_a + K_s \dots \dots \dots (15)$$

where  $K_{cb}$  is a mean crop coefficient at a specific growth stage derived from experimental data where soil moisture is not limiting.  $K_a$  is a coefficient whose value is dependent on the available soil moisture, and  $K_s$  is a coefficient for adjustment of increased evaporation occurring when the soil surface is partially or completely wetted by irrigation or rain (42).

#### Representativity of Meteorological Data

It is well to mention that any data used for the purpose of estimating ET for an irrigated area should be representative of that area. If meteorological data from a nearby desert or dry area are used, for example, from airports or any non-agricultural dry setting for the purpose of estimating ET of an irrigated zone, the corresponding ET values can be misleading. Studies done on this topic (3,7,14) show that introducing irrigation to a previously non-irrigated dry area creates lower day and night time temperatures, higher humidity, and lower windspeeds.

Allen et al.(3) compared the meteorological differences of an irrigated as compared to a non-agricultural setting in an area in southern Idaho. They found that an overestimation of as much as 19% over the season can occur if the weather parameters used for the ETr estimation are not measured in midst of an irrigated area of adequately watered and actively growing crops.

#### ET Estimation for High Altitudes

The difference in vapor pressure between the crop moist surfaces and the surrounding air increases with decrease in the barometric pressure at constant temperature (24). Since changes in barometric pressure are directly related to changes in altitude, all other factors being constant, one would expect an increase in ET with an increase in altitude. Longacre and Blaney (24) observed this effect by measuring pan evaporations at various locations in California ranging in altitude from 150 meters (500 ft) to 2800 meters (9200 ft). It was observed that as the elevation increased the rate of annual evaporation also increased.

Allen and Brockway (2) and Pochop et al.(32) tried to calibrate versions of the Blaney-Criddle ET equation with

temperature as its only measured parameter for high elevations of Idaho. It was found that without an adjustment of 6-10% increase per 1000 m of elevation increase, ET would be underestimated for high elevations. The reason cited was the increased relative cooling of the air during night hours at high altitudes due to decreased density of the atmosphere. This creates low night time temperatures which in turn lower the mean 24 hour air temperature and do not reflect the net effect of daytime temperature and solar radiation available for the ET process.

Johns et al. (23) compared the alfalfa ET estimates obtained from various methods for ten scattered sites in the western U.S.. They found no single equation to function adequately for all ten sites and consequently recommended the usage of correction factors for estimation of field-obtained crop ET. The correction factor found for the Jensen-Haise equation with the elevation correction (JHE) correlated well with changes in altitude. Hence they conclude that out of the seven equations used, "...the JHE equation [with a second correction factor] seems to come closest to being... [a] generally applicable ET method for the sites used" (23, p.20).

Evapotranspiration is just one of several important factors which combine to influence total irrigation water

requirements. Water is lost during the process of delivery to farms and application to fields in the form of evaporation and seepage from canals, laterals, and ditches. Irrigation water is also used for leaching salts, easing tillage or harvest operations, protection against frost, and cooling of plants. Conservation in the use of irrigation water supplies can make available more irrigated agricultural lands. For example, Israel has been able to increase its irrigated lands by 25% without increasing its irrigation water supplies (20). Therefore, if the aim is to conserve water in an arid area, efforts should also be directed toward minimizing losses and increasing efficiencies of use of water other than that of evapotranspiration, otherwise

...there is little reason for insisting upon a method which will give estimates of consumptive use within close limits of accuracy if procedures for estimating the remainder of the water which comprises the total [irrigation] requirements are not of comparable accuracy (33, p.181).

Also the lysimeter should be located at a representative site, i.e. it should be located within a large irrigated alfalfa field.

The measured data available for this study were obtained by Grabow (9) using neutron probes and water-table ly-

## PROCEDURE

Calibration of the Temperature-  
Radiation Equation

Ideally, any calibration of an equation which is to be used as a quantitative estimation of a natural phenomenon, such as evapotranspiration, should be done directly against measured values of that phenomenon. Although sensitive measuring devices, such as various types of lysimeters and neutron probes are available, the use of any of these devices does not necessarily provide the researcher with a reliable longterm data-base. With the main purpose of this study being the calibration of an ETr equation, the subject matter becomes even more complicated since ETr, being an "alfalfa reference evapotranspiration rate" can be hard to measure. In order to measure ETr, a sensitive alfalfa lysimeter would be required with the alfalfa being always at a constant state of full growth and water being non-limiting. Also the lysimeter should be located at a representative site, i.e. it should be located within a large irrigated alfalfa field.

The measured data available for this study were obtained by Grabow (9) using neutron probes and water-table ly-



simeters. Unfortunately these data were not considered as reliable data-base for the calibration of the Temperature-Radiation equation for a number of reasons:

- 1) Often high water tables within the lysimeters created erroneous ET values, particularly with a 7 day interval between readings.
- 2) The lysimeters were not always at a representative setting. For example, a grass lysimeter located at Utah State University's north experimental farm in Logan had only a very small area of grass surrounding it, and a weather station, a tree, and a road were within steps of the lysimeter.
- 3) There were no alfalfa lysimeters in Utah.
- 4) The data were scattered, and often did not cover all the growing season. Where measured data for more than one growing season were available for a site, the ET values for different years were not consistent, that is, the measured ET values fluctuated considerably from one year to another, making them appear unreliable. However, the meteorological data available for this study were generally daily, consistent, continuous (covering all the growing season), and longterm (more than one growing season).

A version of the Penman ET equation (equation 14 of the Literature Review) was calibrated by J. L. Wright at the University of Idaho Agricultural Research Center in Kimberly, Idaho. The data-base used for the Kimberly cali-

bration were derived from several years of measured alfalfa ET using weighing lysimeters at Kimberly (22,42).

Alfalfa has been considered to serve as a good reference crop for a variety of reasons. It is widely grown in arid and semi-arid irrigated areas; 40% of Utah's irrigated acreage is in alfalfa (17). Alfalfa has a long growing season, and produces sufficient canopy thickness in a short period of time. This provides for good absorption of solar radiation above the ground surface. Also, alfalfa has low leaf resistance to water vapor diffusion and has a large root system, especially as compared to grass. Alfalfa produces relatively high ET rates under arid conditions where there are advective sensible heat input available from the air, and its ET rate is little affected by decreasing soil moisture (4,41). The combination of these factors provide the irrigation engineer with a reliable reference crop evapotranspiration rate.

With consideration of the above, the Kimberly Penman method of estimating ETr with the 100 mile wind limit, which has been tested as being an accurate method of alfalfa ETr estimation for Utah (9,17) and Southern Idaho (4,22,40,42), was considered as the best alternative to be used for the calibration of the Temperature-Radiation ETr equation.

The Kimberly Penman ETr equation requires five discrete meteorological measurements. They are: 1) maximum daily temperature  $T_{mx}$ , 2) minimum daily temperature  $T_{mn}$ , 3) early morning dew-point temperature  $T_{dw}$ , 4) incoming daily solar radiation flux  $R_s$ , and 5) daily wind travel. Eighteen sites were selected for which the above data were available. A list of the study sites is shown in Table 2 with their state, county, latitude, longitude, elevation, and years of available data. Among the eighteen sites, there was a total of 45 meteorological datafiles. Figure 5 shows a map of Utah and the relative position of the study sites in the state of Utah.

A version of the FORTRAN program named CRPSIM developed at Utah State University, department of Agricultural and Irrigation Engineering was utilized to read the meteorological datafiles of our study sites and to generate discrete 5-day sum Penman ETr values needed for the calibration process of the Temperature-Radiation equation. For a listing of the CRPSIM program, refer to Appendix II.

The original Jensen-Haise equation had CT and TX values equal to 0.014 and 26.4 respectively. Jensen and Haise estimated these values by simple linear regression (19). In order to come up with coefficients for the Tem-

TABLE 2.- List of the Study Sites with their State, County, Latitude, Longitude, Elevation, and Years of Available Data.

SITE	COUNTY	LATI- TUDE	LONGI- TUDE	ELEVA- TION	YEARS
UTAH					
St. George	Washington	37.08	113.68	2800	1984
Enterprize	Washington	37.57	113.70	5300	1984
Parowon	Iron	37.85	112.83	5930	1980
Flowell	Millard	38.98	112.42	4702	1980,81
Delta	Millard	39.33	112.59	4623	1983
Park City	Summit	40.72	111.52	6740	1982-84
SLC AP	Salt Lake	40.78	111.95	4267	1970-81
Kaysville	Summit	41.07	111.18	4267	1980-82
Paradise	Cache	41.58	111.60	5000	1984
Thornack	Rich	41.75	111.13	6280	1983
Randolph	Rich	41.75	111.13	6280	1982-84
Logan	Cache	41.75	111.82	4580	77,80-82
Garland	Cache	41.75	112.19	4400	1984
IDAHO					
Kimberly	Cassia	42.19	114.12	3960	69-75,80-82
Mont Pelier	Bear Lake	42.32	111.26	5000	1984
Talmage	Caribou	42.70	111.77	5600	1982
WYOMING					
Hilliard Flats	Unita	41.08	111.01	7550	1982-84

FIG. 3.- Map of Utah and the Relative Position of the Study Sites in the State of Utah.

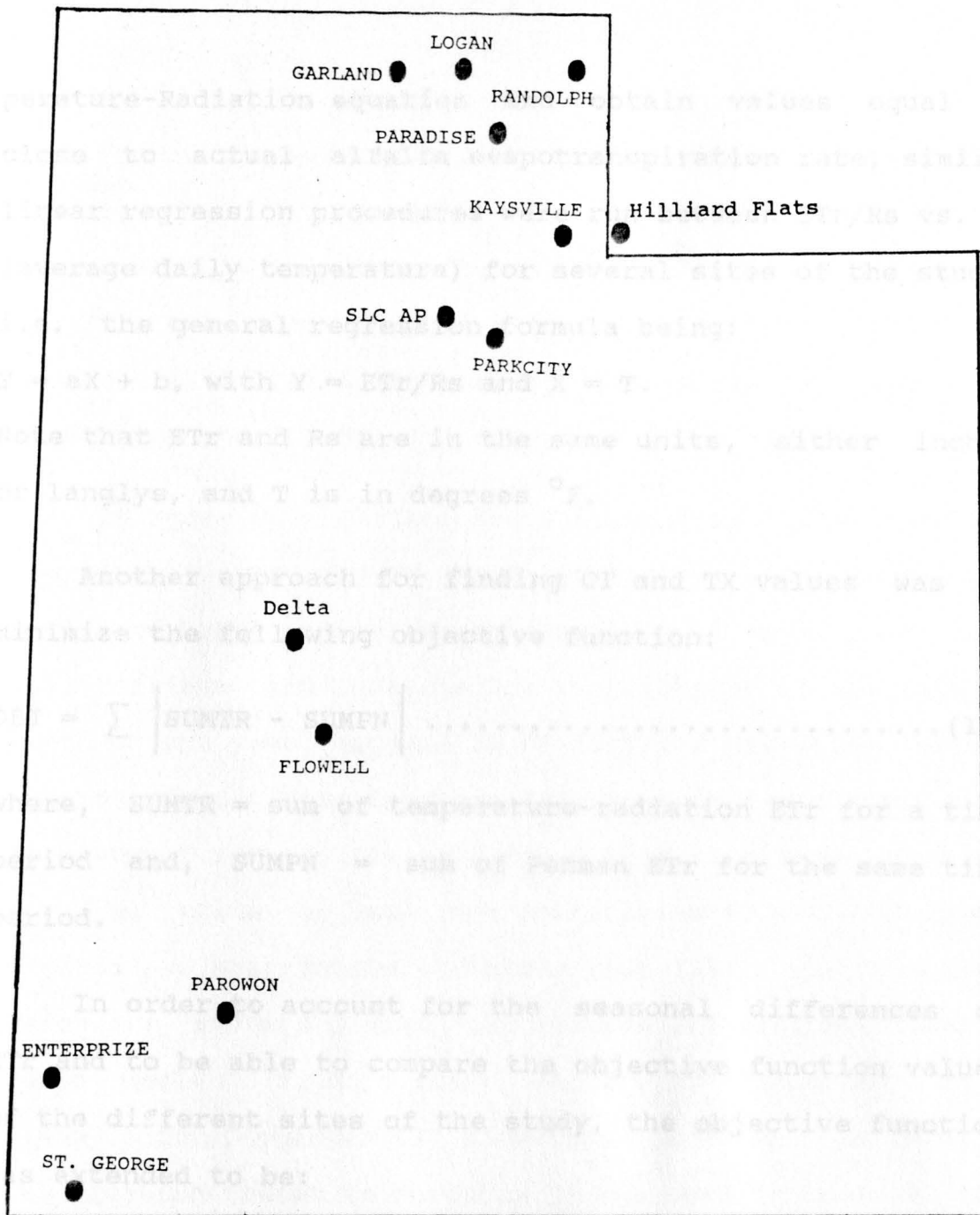


FIG. 5.- Map of Utah and the Relative Position of the Study Sites in the State of Utah.

$$OBJ = \left( \sum |SUMTR - SUMPN| + |SEASTR - SEASPN| \right) \times 100 / SEASPN \dots (17)$$

where, SEASTR = seasonal sum of ETr using the Temperature-radiation method,

SEASPN = seasonal sum of ETr using the Kimberly Penman



perature-Radiation equation and obtain values equal or close to actual alfalfa evapotranspiration rate, similar linear regression procedures were run between  $E_{Tr}/R_s$  vs.  $T$  (average daily temperature) for several sites of the study; i.e. the general regression formula being:  
 $Y = aX + b$ , with  $Y = E_{Tr}/R_s$  and  $X = T$ .  
 Note that  $E_{Tr}$  and  $R_s$  are in the same units, either inches or langlys, and  $T$  is in degrees  $^{\circ}F$ .

Another approach for finding CT and TX values was to minimize the following objective function:

$$OBJ = \sum \left| \text{SUMTR} - \text{SUMPEN} \right| \dots \dots \dots (16)$$

where,  $\text{SUMTR}$  = sum of temperature-radiation  $E_{Tr}$  for a time period and,  $\text{SUMPEN}$  = sum of Penman  $E_{Tr}$  for the same time period.

In order to account for the seasonal differences of  $E_{Tr}$  and to be able to compare the objective function values of the different sites of the study, the objective function was extended to be:

$$OBJ = \left( \sum \left| \text{SUMTR} - \text{SUMPEN} \right| + \left| \text{SEASTR} - \text{SEASPN} \right| \right) \times 100 / \text{SEASPN} \dots \dots (17)$$

where,  $\text{SEASTR}$  = seasonal sum of  $E_{Tr}$  using the Temperature-Radiation method,

$\text{SEASPN}$  = seasonal sum of  $E_{Tr}$  using the Kimberly Penman

method,

OBJ is in units of inches/inches and represented as a % value.

A FORTRAN computer program, named TEMPRAD, was written to find the lowest objective function for every data file. Appendix II contains the FORTRAN program listings used for this study and a sample of a typical data-file used.

In order to calculate the Jensen-Haise 1970 equation coefficients CT and TX, the elevation of the site and the saturated vapor pressures corresponding to the long term mean maximum and mean minimum temperatures of the warmest month of the year (assumed to be July for the study sites) were required. Mean longterm July maximum and minimum temperatures for the sites were found from isolines outlined for the state of Utah in a publication from the National Oceanic and Atmospheric Administration (25). The isolines were the average normal July maximum and minimum temperatures derived from twenty years of recorded data (1932-1952). Appendix III shows such isolines. An empirical equation (refer to Appendix I) was used to derive the saturated vapor pressure as a function of temperature. Values for the coefficients CT and TX were obtained for the Jensen-Haise 1970 equation using eq.s 6b-6f of Chapter 1. Table 3 contains the mean long term July maximum and mini-

mum temperatures, and the CT and TX coefficients for all the sites in this study.

When trying to find the lowest OBJ function for the Temperature-Radiation equation, TEMPRAD went through a simple do-loop, each time using a distinct pair of CT and TX. Figure 6 shows a 3-dimensional graph with changes of CT and TX, and the corresponding change in the objective function OBJ for Logan 1982 growing season's data.

TEMPRAD was modified to find the CT values for each site while holding TX equal to zero. Holding TX as zero transferred the temperature-radiation equation into a version of the Hargreaves equation. It was found that by doing so, the OBJ values did not increase considerably. The final calibration results and the corresponding coefficients to be used in the Datapod field computer are stated in the Results section.

#### General Method for Determining CT

Equations 6b-6f of the Literature Review section were proposed by Jensen et al.(21) for determining CT and TX of the Jensen-Haise equation. Jensen et al. took CT and TX as functions of elevation and saturation vapor pressure de-

TABLE 3.- Mean Long Term July Maximum and Minimum Temperatures, and the CT and TX Coefficients of the 1970 Jensen-Haise ET Equation.

SITE	July long term average temperatures (F)		CT	TX
	Max.	Min.		
St. George	101.0	68.6	0.0138	13.60
Enterprise	90.0	57.2	0.0145	14.08
Parowon	88.0	56.0	0.0147	14.03
Flowell	90.0	60.0	0.0136	15.10
Delta	94.0	60.0	0.0145	13.55
Park City	79.4	41.4	0.0144	14.39
Kaysville	92.3	60.8	0.0139	14.75
SLC AP	92.3	60.8	0.0139	14.75
Paradise	83.0	53.4	0.0131	16.31
Thornack	80.6	43.8	0.0142	14.72
Randolph	81.1	44.2	0.0143	14.61
Logan	87.5	56.0	0.0136	15.56
Garland	87.3	58.4	0.0132	16.16
Kimberly	84.0	54.0	0.0123	17.11
Mont Pelier	81.9	45.8	0.0135	15.80
Talmage	84.0	49.0	0.0141	14.87
Hilliard flts	74.7	45.3	0.0134	15.18

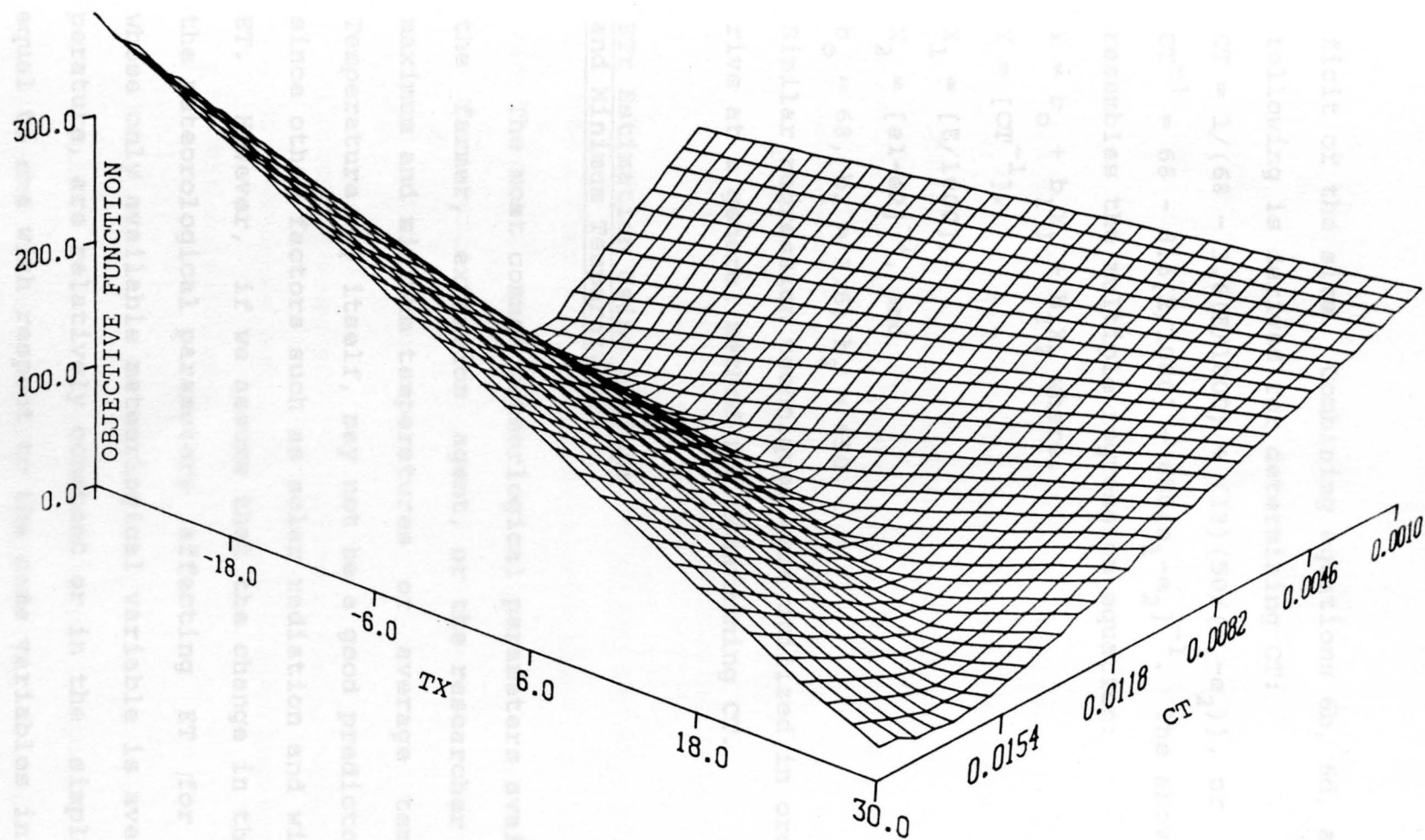


FIG. 6- Three Dimensional Graph Showing the Objective Function with Changes in CT and TX Coefficients Using the Logan 1982 Meteorological Data.



ficit of the site. Combining equations 6b, 6d, and 6e, the following is derived for determining CT:

$$CT = 1/(68 - 3.6(E/1000) + (13)(50/(e_1 - e_2))), \text{ or}$$

$CT^{-1} = 68 - 3.6[E/1000] + 650[e_1 - e_2]^{-1}$ . The above equation resembles the multiple regression equation:

$$Y = b_0 + b_1X_1 + b_2X_2 \text{ where,}$$

$$Y = [CT^{-1}],$$

$$X_1 = [E/1000],$$

$$X_2 = [e_1 - e_2]^{-1}, \text{ and}$$

$$b_0 = 68, b_1 = 3.6, b_2 = 650.$$

Similar regression techniques were utilized in order to arrive at a general method for determining CT.

#### ETr Estimation Using Maximum and Minimum Temperatures

The most common meteorological parameters available to the farmer, extension agent, or the researcher are daily maximum and minimum temperatures or average temperature. Temperature, by itself, may not be a good predictor of ETr, since other factors such as solar radiation and wind affect ET. However, if we assume that the change in the rest of the meteorological parameters affecting ET for site A, whose only available meteorological variable is average temperature, are relatively constant or in the simplest case equal to one with respect to the same variables in a nearby

site B, that is to say, if we assume that site A and B's weather parameters, i.e. vapor pressure, dew point, solar radiation, wind,...etc. are the same except for daily maximum and minimum temperatures, theoretically we could estimate ETr for site A by using ETr values of site B.

In order to see what effect temperature change alone has on ETr, TEMPRAD was modified to calculate Penman ETr while changing maximum and minimum temperatures within the range of  $\pm 10^{\circ}\text{F}$ . It was found that if all other factors are held constant, temperature change in the Penman equation would be linearly related to ETr. Figure 8 of the next chapter shows this linear relationship.

Assuming Kaysville is the site B we were referring to, ETr for a nearby site A can be equivalent to the product of ETr of site B,  $\text{ETr}_B$ , and a multiplier "m". m would simply be the ratio of  $\text{ETr}_B$  with a temperature change to  $\text{ETr}_B$  without a temperature change.

Attempts were made to derive a single equation relating the average temperature T, and the temperature change  $\Delta T$  to the multiplier m.

As stated in the Literature Review section, Hargreaves and Samani (13) proposed an ET equation with the required measurements of maximum and minimum temperatures, and esti-

mated extraterrestrial radiation. The equation is:

$$ET_0 = K \times ( T \times TD^{0.05} \times Ra ) \quad \text{where,}$$

$ET_0$  = grass reference evapotranspiration rate (in/time period),

$Ra$  = extraterrestrial radiation (same units as  $ET_0$ ),

$TD$  = difference of mean maximum and mean minimum temperatures ( $^{\circ}F$ ).

The above equation was calibrated against eight years of Alta fescue grass ET from lysimeters at Davis-CA (13). The K found was 0.00094. Hargreaves and Samani suggested this K for places where local calibration is not possible. Assuming a 1.2 factor of conversion between grass ET and alfalfa ET, the suggested K for the reference crop of alfalfa would be 0.001128 (1.2 x 0.00094).

In order to find locally calibrated K values for the sites of this study, TEMPRAD was used to minimize the same objective function, OBJ, as was utilized for the Temperature-Radiation equation. These and other results are stated in the next chapter.

TABLE 4.- Results of the Regression Between ETRs and Average Temperature, T, for Several Datafiles.

## RESULTS

### Calibration of the Temperature-Radiation Equation

Table 4 shows the linear regressions between ETR/Rs and T for several years of growing season's data from Flowell, Kaysville, Kimberly, and Logan-Utah. As can be seen from the low correlation coefficients  $R^2$ , such linear regression, a technique used by Jensen and Haise (19) was not a good strategy for the calibration of the Temperature-Radiation equation with the meteorological data available for the sites of this study.

Seasonal data were reduced to the months of May through September wherever possible. The objective function as stated by equation 17 was used in the TEMPRAD program. TEMPRAD was run with the datafiles available. The values for CT and TX were restricted to 0.005 to 0.020, and -30 to +30 respectively. Consequently the objective functions were minimized and distinct pairs of CT and TX were derived for each datafile. The initial calibration of the Temperature-Radiation method against the Penman method are shown in Table 5. Table 5 contains all the study sites and

TABLE 4.- Results of the Regression Between ET/RS and Average Temperature, T, for Several Datafiles.

SITE	YEAR	PERIOD	REGRESSION EQUATION	R Squared
FLOWELL	1980	5/01-09/30	(ET/RS)=.0099 + .0672(T)	45.0%
FLOWELL	1981	4/08-10/31	(ET/RS)=.0100 + .2527(T)	10.0%
KAYSVILE	1980	4/01-10/03	(ET/RS)=.0055 + .3900(T)	5.4%
KAYSVILE	1981	4/01-10/04	(ET/RS)=.0050 + .3959(T)	3.8%
KAYSVILE	1982	4/06-10/20	(ET/RS)=.0069 + .2369(T)	19.7%
KIMBERLY	1980	4/01-09/30	(ET/RS)=.0076 + .2122(T)	7.8%
KIMBERLY	1981	4/01-09/30	(ET/RS)=.0046 + .4617(T)	4.9%
KIMBERLY	1982	3/16-09/30	(ET/RS)=.0053 + .2704(T)	14.8%
LOGAN	1981	4/01-10/30	(ET/RS)=.0055 + .1915(T)	10.9%
LOGAN	1982	4/01-09/30	(ET/RS)=.0041 + .3138(T)	5.4%



the initial CT, TX, and OBJ values obtained. As can be seen no site produced the same CT and TX coefficients for all years of data available for that site.

In order to avoid negative TX values, and to attempt to reduce the coefficients pertaining to each site from two (CT and TX) to one (CT), TEMPRAD was modified to find the CT coefficient while holding TX at zero. By doing this, CT values of the different years for the same site tended to take the same value. Table 6 shows the CT coefficients when TX is held at zero and the corresponding OBJ functions obtained. The OBJ functions of Table 6 are on the average only 2.25% higher than their corresponding values in Table 5.

For sites for which only one growing season's meteorological data were available, the CT values obtained were considered as final and are recommended as the coefficients to be used by the Datapod field computer. However, for sites which had more than one growing season's data available, and different CT coefficients were obtained corresponding to each year's data, a method was utilized to find the best CT to be used by the Datapod at those sites.

TABLE 5.-Initial Calibration Results of the Temperature-Radiation Equation with Distinct Pairs of CT and TX for Each Datafile.

SITE	YEAR	GROWING SEASON	CT	TX	OBJ
St. George	1984	5/ 1- 9/30	0.0080	-25.0	5.62%
Enterprize	1984	5/ 1- 9/30	0.0075	-30.0	4.19%
Parowon	1980	5/ 1- 9/30	0.0095	- 1.0	3.65%
Flowell	1980	5/ 1- 9/30	0.0130	12.0	4.74%
Flowell	1981	5/ 1- 9/30	0.0080	-29.0	11.83%
Delta	1983	5/ 1- 9/30	0.0100	5.0	4.63%
Park City	1982	5/27-10/ 9	0.0070	-27.5	5.17%
Park City	1983	5/ 5-10/ 5	0.0070	-24.5	5.56%
Park City	1984	5/ 4- 9/30	0.0070	-29.5	4.44%
SLC AP	1972	5/ 1- 9/30	0.0090	2.0	5.53%
SLC AP	1973	5/ 1- 9/30	0.0080	- 7.5	5.56%
SLC AP	1974	5/ 1- 9/30	0.0070	-19.0	3.93%
SLC AP	1975	5/ 1- 9/30	0.0085	- 1.5	4.20%
SLC AP	1976	5/ 1- 9/30	0.0075	-11.0	3.92%
SLC AP	1977	5/ 1- 9/30	0.0070	-17.0	6.04%
SLC AP	1978	5/ 1- 9/30	0.0085	- 2.5	5.36%
SLC AP	1979	5/ 1- 9/30	0.0065	-24.0	4.47%
SLC AP	1980	5/ 1- 9/30	0.0095	6.0	5.53%
SLC AP	1981	5/ 1- 9/30	0.0070	-19.0	6.39%
Kaysville	1980	5/ 1- 9/30	0.0115	9.0	6.88%
Kaysville	1981	5/ 1- 9/30	0.0065	-28.5	6.52%
Kaysville	1982	5/ 1- 9/30	0.0075	-21.0	5.92%
Logan	1980	5/ 1- 9/30	0.0075	-13.5	5.54%
Logan	1981	5/ 1- 9/30	0.0070	-20.5	4.95%
Logan	1982	5/ 1- 9/30	0.0085	- 8.0	6.29%
Paradise	1984	5/ 1- 9/30	0.0070	-29.5	4.88%
Thornack	1983	5/ 1- 9/30	0.0120	8.0	5.75%
Randolph	1982	5/26-10/22	0.0080	-15.5	5.42%
Randolph	1983	5/ 1- 9/30	0.0065	-28.5	4.02%
Randolph	1984	5/ 1- 9/30	0.0070	-26.5	4.58%

Garland	1984	5/ 1- 9/30	0.0075	-21.0	5.18%
Kimberly	1969	5/ 1- 9/30	0.0105	3.5	5.47%
Kimberly	1970	5/ 1- 9/30	0.0080	-13.5	5.32%
Kimberly	1972	5/ 1- 9/30	0.0095	- 2.5	5.82%
Kimberly	1973	5/ 1- 9/30	0.0075	-18.5	4.40%
Kimberly	1974	5/ 1- 9/30	0.0085	- 9.0	4.49%
Kimberly	1975	5/ 1- 9/30	0.0080	-13.0	5.05%
Kimberly	1980	5/ 1- 9/30	0.0100	0.5	8.65%
Kimberly	1981	5/ 1- 9/30	0.0080	-14.5	5.16%
Kimberly	1982	5/ 1- 9/30	0.0065	-23.5	5.04%
Mont Pelier	1984	5/ 1- 9/30	0.0065	-28.5	5.81%
Talmage	1982	5/ 7-10/ 7	0.0065	-26.0	6.35%
Hilliard Flts	1982	5/27-10/22	0.0075	-18.5	6.03%
Hilliard Flts	1983	5/ 6-10/ 6	0.0080	-15.0	5.06%
Hilliard Flts	1984	5/ 4-10/ 2	0.0070	-29.0	4.71%
				Average OBJ:	5.42%

TABLE 6.- CT Coefficients when TX is Held at Zero and the Corresponding OBJ Function Values Obtained.

SITE	YEAR	TX=0 CT	OBJ
St. George	1984	0.01054	6.40%
Enterprize	1984	0.01092	6.22%
Parowon	1980	0.00966	5.12%
Flowell	1980	0.01069	7.00%
Flowell	1981	0.01131	15.42%
Delta	1983	0.00923	7.50%
Park City	1982	0.01045	7.41%
Park City	1983	0.01016	8.83%
Park City	1984	0.01081	8.48%
SLC AP	1972	0.00874	8.37%
SLC AP	1973	0.00886	7.67%
SLC AP	1974	0.00889	5.44%
SLC AP	1975	0.00868	6.71%
SLC AP	1976	0.00867	6.48%
SLC AP	1977	0.00869	8.58%
SLC AP	1978	0.00879	8.45%
SLC AP	1979	0.00868	6.73%
SLC AP	1980	0.00865	7.19%
SLC AP	1981	0.00887	8.30%
Kaysville	1980	0.00997	7.04%
Kaysville	1981	0.00922	9.01%
Kaysville	1982	0.00986	7.95%
Logan	1980	0.00953	6.07%
Logan	1981	0.00916	6.95%
Logan	1982	0.00906	6.62%
Paradise	1984	0.01026	8.01%
Thornack	1983	0.01035	7.31%
Randolph	1982	0.01017	8.13%
Randolph	1983	0.00982	8.60%
Randolph	1984	0.01040	7.99%
Garland	1984	0.00992	7.20%
Kimberly	1969	0.00994	6.11%
Kimberly	1970	0.00974	8.00%

Kimberly	1972	0.00999	7.51%
Kimberly	1973	0.00970	6.74%
Kimberly	1974	0.00970	6.58%
Kimberly	1975	0.00969	7.25%
Kimberly	1980	0.00992	9.46%
Kimberly	1981	0.00983	7.40%
Kimberly	1982	0.00898	6.67%
Mont Pelier	1984	0.00971	9.86%
Talmage	1982	0.00950	7.04%
Hilliard Flts	1982	0.01013	8.55%
Hilliard Flts	1983	0.01026	8.35%
Hilliard Flts	1984	0.01077	8.50%

Average OBJ: 7.67%

Table 7 contains all the study sites with the final CT coefficients that were recommended to be used by the Datapod field computer. The CT values can be round off to four decimal digits to match the Datapod specifications.

The Temperature-Radiation equation which was calibrated here is actually a version of the Jensen-Haise equation. It was found that out of the 45 datafiles used, 42 produced a smaller OBJ function corresponding to the final CT coefficient as compared to the OBJ obtained when using the 1970 Jensen-Haise coefficients.



If the OBJ function of a site be plotted against the changes of the CT coefficient, a "V" shaped graph will be created for every year. The method used for finding a single CT for a site was to sum up the OBJ values corresponding to distinct CT values for the years of data available for that site. The CT corresponding to the lowest sum of OBJs was taken as the Datapod's CT coefficient for that site. This is the same as drawing an average-fit graph. The lowest point of the graph represents the final CT coefficient for that site. Figure 7 shows the average of the three year OBJ sum as plotted against CT and compares it to the Hilliard Flats' 1982, 83, and 84 OBJ functions.

Table 7 contains all the study sites with the final CT coefficients that were recommended to be used by the Datapod field computer. The CT values can be round off to four decimal digits to match the Datapod specifications.

The Temperature-Radiation equation which was calibrated here is actually a version of the Jensen-Haise equation. It was found that out of the 45 datafiles used, 42 produced a smaller OBJ function corresponding to the final CT coefficient as compared to the OBJ obtained when using the 1970 Jensen-Haise coefficients.

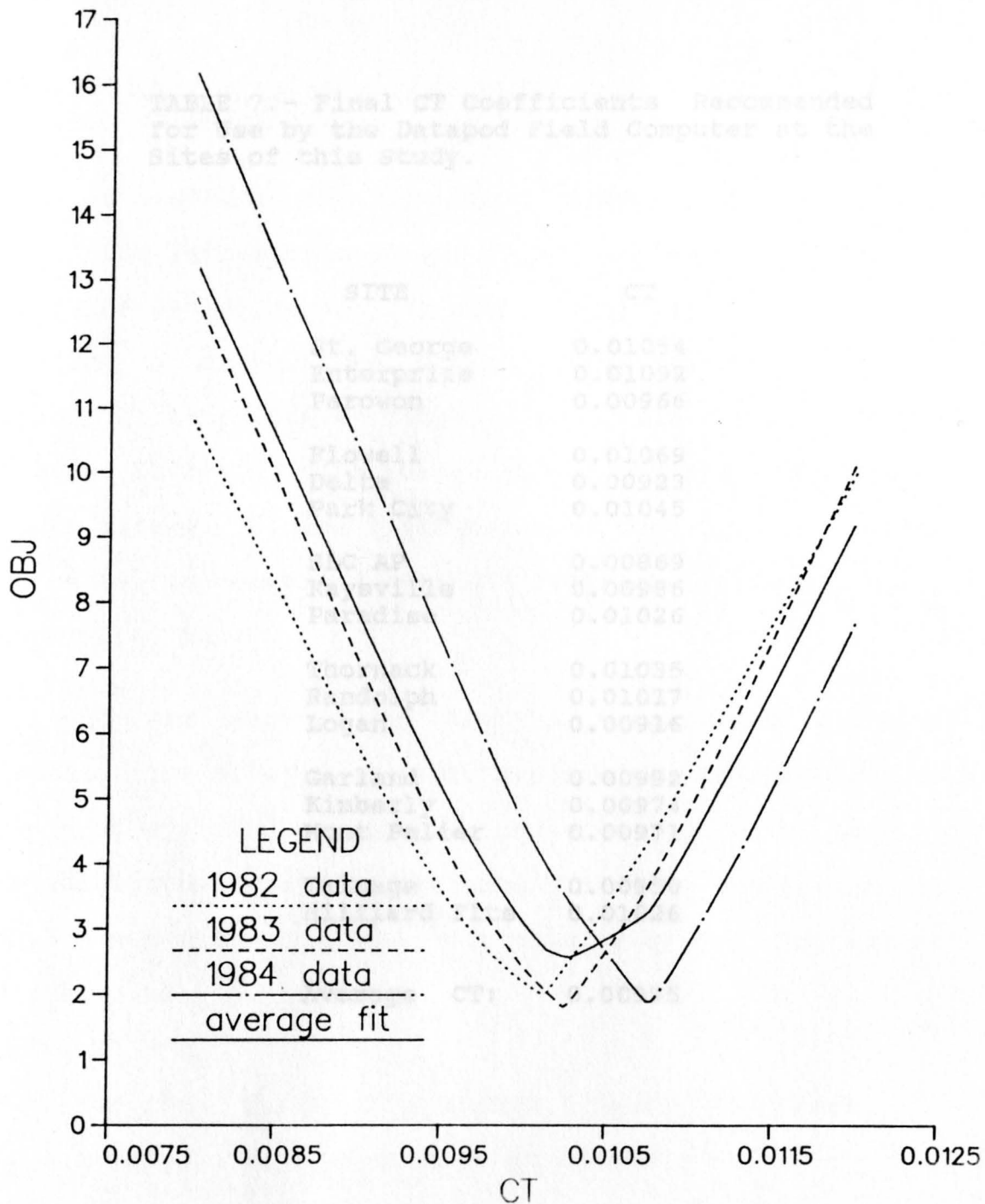


FIG. 7.- Hilliard Flats, Wyoming's 1982, 83 and 84 Objective Functions as Plotted Against a Range of CT Values, and the Average-Fit Curve Obtained.

TABLE 7.- Final CT Coefficients Recommended for Use by the Datapod Field Computer at the Sites of this Study.

SITE	CT
St. George	0.01054
Enterprise	0.01092
Parowon	0.00966
Flowell	0.01069
Delta	0.00923
Park City	0.01045
SLC AP	0.00869
Kaysville	0.00986
Paradise	0.01026
Thornack	0.01035
Randolph	0.01017
Logan	0.00916
Garland	0.00992
Kimberly	0.00974
Mont Pelier	0.00971
Talmage	0.00950
Hilliard Flts	0.01026
Average CT:	0.00995

$$\Delta T = T_A - T_B$$

$$a = [ETR_B(T + \Delta T) / ETR_B(T)] = [ETR_A(T) / ETR_B(T)] \text{ where,}$$

$T_A$  = average daily temperature at site A ( $^{\circ}$ F),

$T_B$  = average daily temperature at site B ( $^{\circ}$ F),

$ETR_B(T)$  = daily ETR at site B with the actual temperature of that day,

$ETR_B(T + \Delta T)$  = daily ETR rate at site B with theoretical temperature  $T + \Delta T$ .  $ETR_B(T + \Delta T)$  is assumed to be equiva-

ETr Estimation Using Maximum and Minimum Temperatures

As stated in the Procedure, TEMPRAD was modified to vary the temperature in the Penman ETr equation within  $\pm 10$   $^{\circ}\text{F}$ . The ratio between ETr with a temperature change to ETr without a temperature change was designated as "m". For specific temperature ranges, the relationship between change in the average temperature,  $\Delta T$ , and the multiplier m was linear. Figure 8 demonstrates such linear relationships for the combined 1981 and 1982 meteorological data of Kaysville, UT.

$\Delta T$  was assumed to be the temperature difference between site A, with the only available meteorological variable of average temperature, and site B, with available meteorological data to estimate ETr using the Penman method. m was assumed as the multiplier required to estimate ETr at site A, that is:

$$\Delta T = T_A - T_B$$

$$m = [\text{ETr}_B(T + \Delta T) / \text{ETr}_B(T)] \approx [\text{ETr}_A(T) / \text{ETr}_B(T)] \text{ where,}$$

$T_A$  = average daily temperature at site A ( $^{\circ}\text{F}$ ),

$T_B$  = average daily temperature at site B ( $^{\circ}\text{F}$ ),

$\text{ETr}_B(T)$  = daily ETr at site B with the actual temperature of that day,

$\text{ETr}_B(T + \Delta T)$  = daily ETr rate at site B with theoretical temperature  $T + \Delta T$ .  $\text{ETr}_B(T + \Delta T)$  is assumed to be equiva-

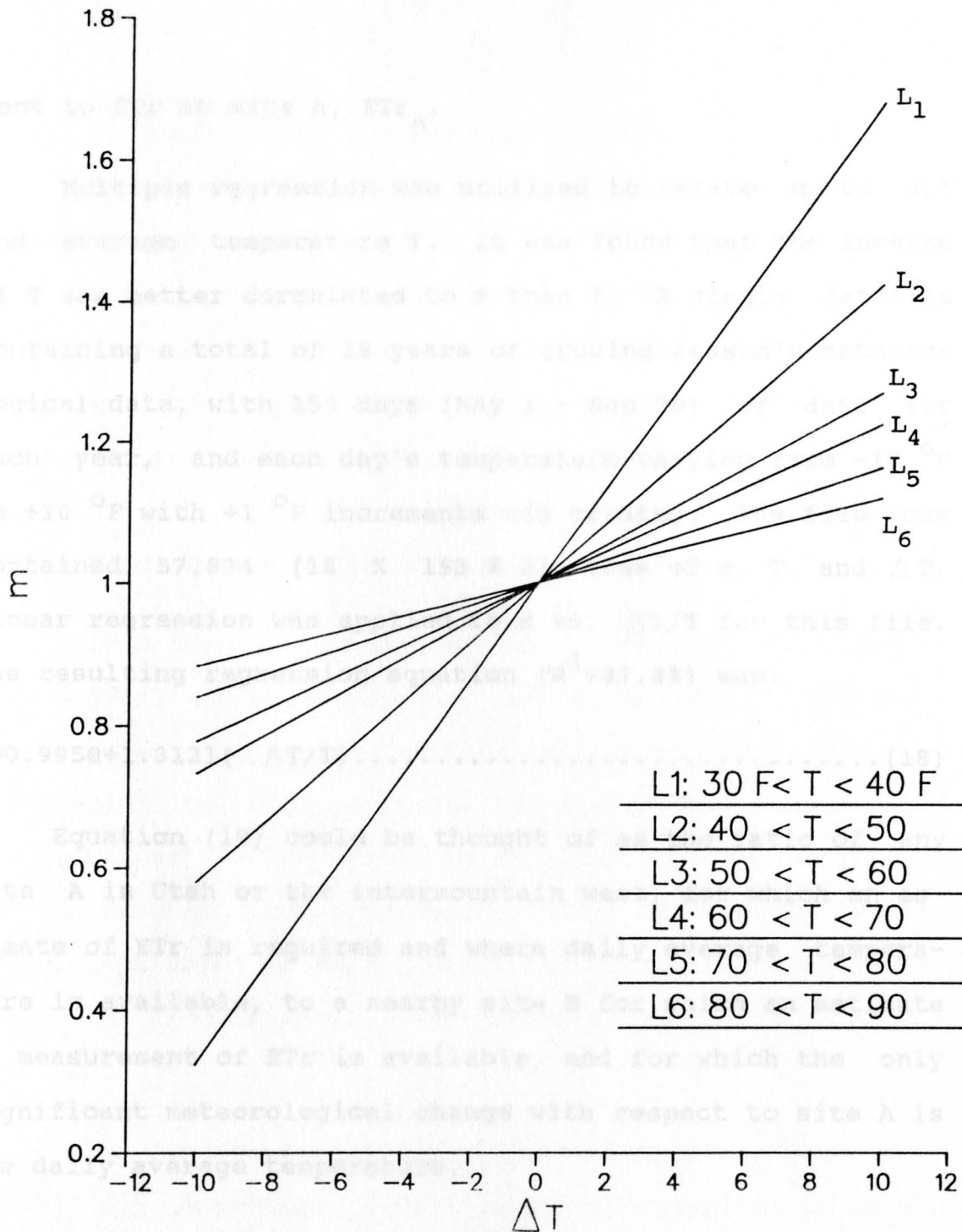


FIG. 8.- Linear Regression Lines Between Multiplier "m", and Change in Temperature  $\Delta T$ , for Specific Temperature Ranges of Kaysville 1981, and 1982 Meteorological Data (Penman ETr Used).



lent to ETr at site A,  $ETr_A$ .

Multiple regression was utilized to relate  $m$  to  $\Delta T$  and average temperature  $T$ . It was found that the inverse of  $T$  was better correlated to  $m$  than  $T$ . A single datafile containing a total of 18 years of growing season's meteorological data, with 153 days (May 1 - Sep 30) of data for each year, and each day's temperature varying from  $-10^{\circ}F$  to  $+10^{\circ}F$  with  $+1^{\circ}F$  increments was created. The file thus contained 57,834 ( $18 \times 153 \times 21$ ) rows of  $m$ ,  $T$ , and  $\Delta T$ . Linear regression was applied to  $m$  vs.  $\Delta T/T$  for this file. The resulting regression equation ( $R^2=87.8\%$ ) was:

$$m=0.9958+1.3121(\Delta T/T) \dots \dots \dots (18)$$

Equation (18) could be thought of as the ratio of any site A in Utah or the intermountain west, for which an estimate of ETr is required and where daily average temperature is available, to a nearby site B for which an estimate or measurement of ETr is available, and for which the only significant meteorological change with respect to site A is the daily average temperature.

#### Calibration of the New Hargreaves Equation

Another way to estimate ETr using temperature data alone is by an ETr equation with the only required measure-

ments of maximum and minimum temperatures. As stated in the Procedure, the New Hargreaves equation was calibrated using TEMPRAD and the same objective function as was for the Temperature-Radiation equation. Table 8 shows the initial calibration results and the minimum OBJ functions achieved for the New Hargreaves equation. The same procedure was used to arrive at a single coefficient for each site as was used for the Temperature-Radiation equation. Table 9 presents the final calibration results.

Since the New Hargreaves equation requires the extra-terrestrial solar radiation,  $R_a$ , as an input, Appendix V was created. Appendix V contains a list of  $R_a$  values corresponding to Utah latitudes for the growing season April 1- September 30. The formulas required for the estimation of  $R_a$  were acquired from Andrew Keller, Utah State University, department of Agricultural and Irrigation Engineering. The program used in estimating  $R_a$  is listed in Appendix II.

Five-day ETr values for all the study sites were combined and datafiles were created corresponding to Penman ETr, Jensen-Haise 1970 (J-H 70) ETr, Temperature-Radiation (T-R) ETr, New Hargreaves with fixed K (Har-fK) ETr, and Calibrated New Hargreaves (Har-calk) ETr equations. The four equations stated were regressed against the Penman

TABLE 8.-Initial Calibration Results of the New Hargreaves ETr Equation with K Coefficients and Minimum Objective Functions Found.

SITE	YEAR	K	OBJ
St. George	1984	0.001057	5.69%
Enterprize	1984	0.001133	7.34%
Parowon	1980	0.001161	5.37%
Flowell	1980	0.000998	4.99%
Flowell	1981	0.000937	16.47%
Delta	1983	0.001069	7.46%
Park City	1982	0.001009	8.21%
Park City	1983	0.001038	9.22%
Park City	1984	0.001065	9.56%
SLC AP	1972	0.001040	6.27%
SLC AP	1973	0.001011	6.20%
SLC AP	1974	0.000999	4.36%
SLC AP	1975	0.001067	3.81%
SLC AP	1976	0.001042	4.91%
SLC AP	1977	0.001035	4.12%
SLC AP	1978	0.001031	6.25%
SLC AP	1979	0.001032	5.27%
SLC AP	1980	0.001040	5.12%
SLC AP	1981	0.001010	5.48%
Kaysville	1980	0.001076	7.88%
Kaysville	1981	0.000965	6.43%
Kaysville	1982	0.001136	5.92%
Logan	1980	0.001078	7.69%
Logan	1981	0.001094	8.21%
Logan	1982	0.001124	8.37%
Paradise	1984	0.001073	8.06%
Thornack	1983	0.001080	8.83%
Randolph	1982	0.001062	6.60%
Randolph	1983	0.001074	8.01%
Randolph	1984	0.001085	8.26%
Garland	1984	0.001081	11.99%
Kimberly	1969	0.001146	5.31%
Kimberly	1970	0.001116	5.86%

Kimberly	1972	0.001132	5.57%
Kimberly	1973	0.001155	6.10%
Kimberly	1974	0.001189	7.38%
Kimberly	1980	0.001093	6.69%
Kimberly	1981	0.001143	5.77%
Kimberly	1982	0.001017	8.67%
Mont Pelier	1984	0.001018	8.54%
Talmage	1982	0.000983	7.30%
Hilliard Flts	1982	0.001147	6.19%
Hilliard Flts	1983	0.001122	7.97%
Hilliard Flts	1984	0.001173	7.91%

Park City

0.001038

SAC AP

Average OBJ: 7.08%

Kaysville

0.001077

Logan

0.001094

Paradise

0.001073

Thornack

0.001080

Randolph

0.001074

Garland

0.001081

Kimberly

0.001135

Mont Pelier

0.001018

Talmage

0.000983

Hilliard Flts

0.001147

Average K:

0.001073

TABLE 9.- Final Calibration Results of  
New Hargreaves ETr Equation for the  
Sites of this Study.

SITE	K
St. George	0.001057
Enterprize	0.001133
Parowon	0.001161
Flowell	0.000998
Delta	0.001069
Park City	0.001038
SLC AP	0.001032
Kaysville	0.001075
Logan	0.001094
Paradise	0.001073
Thornack	0.001080
Randolph	0.001074
Garland	0.001081
Kimberly	0.001135
Mont Pelier	0.001018
Talmage	0.000983
Hilliard Flts	0.001147
Average K:	0.001073



method. Table 10 shows the regression equations and the correlation coefficients,  $R^2$ , achieved.

In order to demonstrate how well the four equations mentioned estimate ETr, and whether or not they will underestimate or overestimate the evapotranspiration irrigation requirements, the following formula was utilized for comparison of the 5-day ETr estimate datafiles:

$$\%d = [(SUM5 - SUM5PN)/SUM5PN] \times 100 \dots\dots\dots(19)$$

SUM5=5 day sum of ETr using an ETr method, SUM5PN=5 day sum of ETr using the Penman ETr method, and %d=percentage of underestimation (if negative), or overestimation (if positive).

Table 11 shows the percentage of 5-day estimates by each of the four methods lying within 5%, 10%, and 15% of the Penman ETr method. Figures 9 through 12 are plots of %d vs. the frequency of occurrence corresponding to the J-H 70, T-R, Har-fk, Har-fcal ETr equations respectively. As can be seen the Temperature-Radiation equation ranks as the best estimator among the four methods with 90.9% of the 5-day ETr estimates within  $\pm 15\%$  of Penman, followed by the calibrated New Hargreaves, Har-fK, and J-H 70 ETr equations each with 89.3%, 79.7%, and 55.1% of estimates within  $\pm 15\%$  of Penman estimates respectively.

TABLE 10.- Regression Equations Obtained by Running 5-Day Discrete ETr Estimates Found by Jensen-Haise 1970, Temperature-Radiation, New Hargreaves, and Calibrated New Hargreaves ETr Methods Against the 5-Day Penman ETr Estimates for all Sites Combined.

	Regression Equation	$R^2$
Jensen-Haise	$ETr = -0.2290 + 1.280(\text{Penman ETr})$	83.6 %
Temp.-Rad.	$ETr = -0.0746 + 1.060(\text{Penman ETr})$	93.4 %
New Harg.	$ETr = 0.0684 + 0.995(\text{Penman ETr})$	85.6 %
Calib. New Harg.	$ETr = 0.0533 + 0.953(\text{Penman ETr})$	89.2 %

TABLE 11.-Percentage of 5-Day ETr Estimates Using the Temperature-Radiation, Jensen-Haise 1970, New Hargreaves, and Calibrated New Hargreaves ETr Equations within 5%, 10%, and 15% of the Penman ETr Method for all Sites Combined.

ETr Equation	Percentage of 5-day data within		
	5%	10%	15% of Penman
Temp.-Radiation	46.6%	78.3%	90.9%
New Har. calib.	42.1%	73.3%	89.3%
New Hargreaves	32.8%	59.4%	79.7%
Jensen-Haise	23.6%	41.1%	55.1%

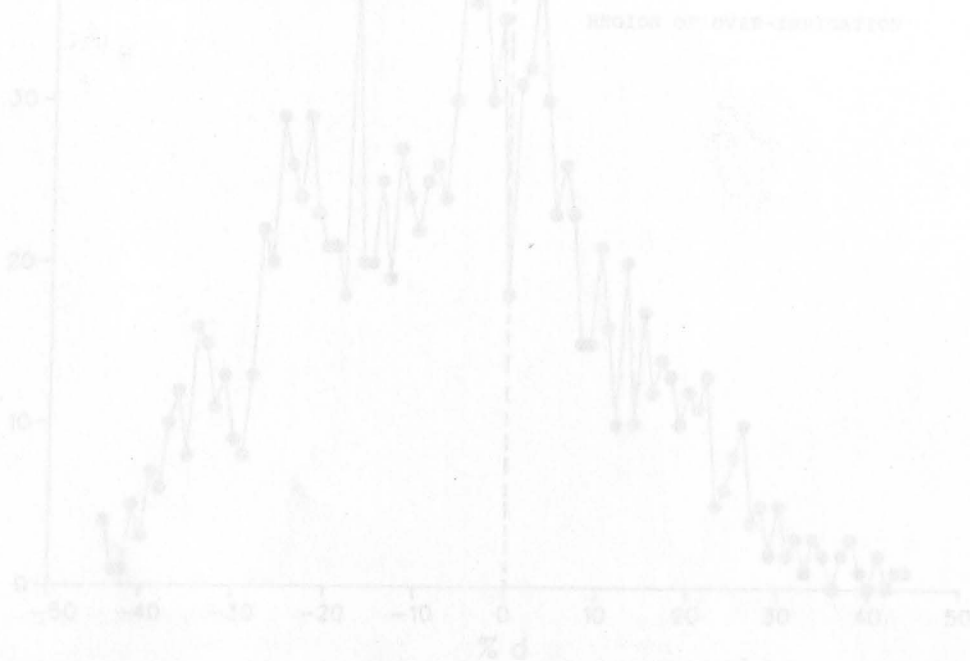


FIG. 9.- Plot of %d vs. Frequency of Occurrence for the Jensen-Haise 1970 ETr Equation as Compared to Penman ETr (All Sites Combined).

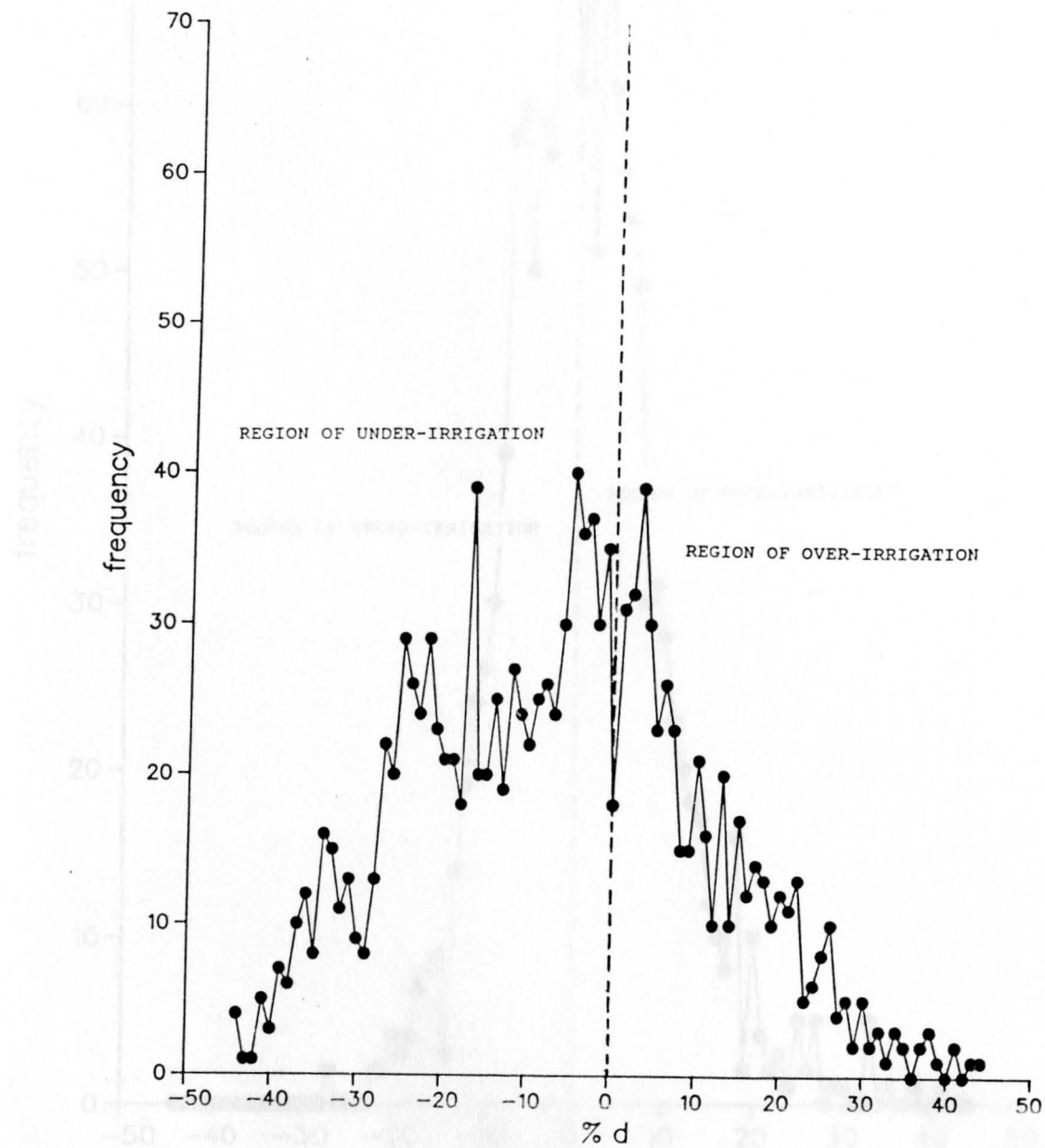


FIG. 9.- Plot of %d vs. Frequency of Occurrence for the Jensen-Haise 1970 ETr Equation as Compared to Penman ETr (All Sites Combined).

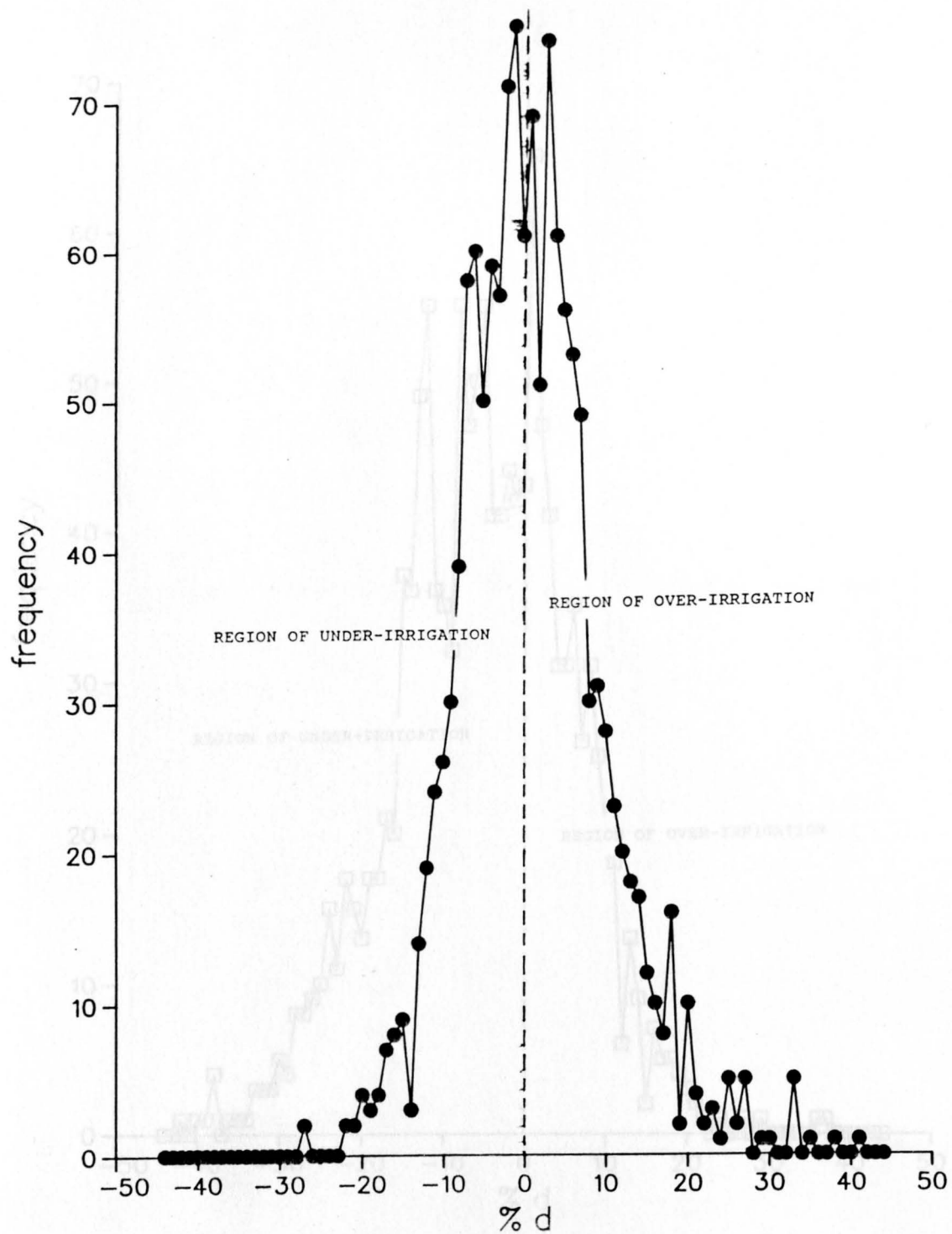


FIG. 10.- Plot of %d vs. Frequency of Occurrence for the Temperature-Radiation ETr Equation as Compared to Penman ETr (All Sites Combined).



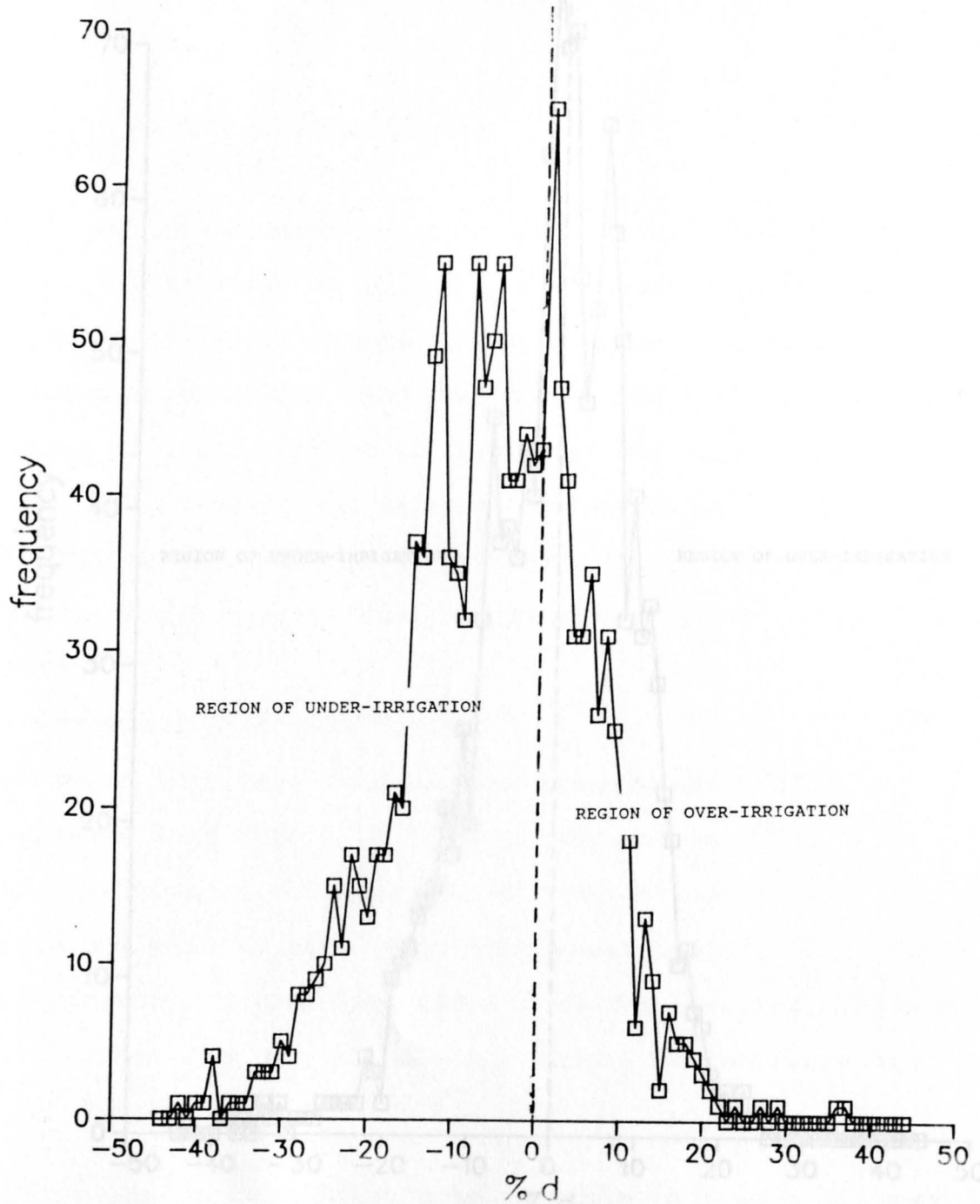


FIG. 11.- Plot of %d vs. Frequency of Occurrence for the New Hargreaves ETR Equation as Compared to Penman ETR (All Sites Combined).

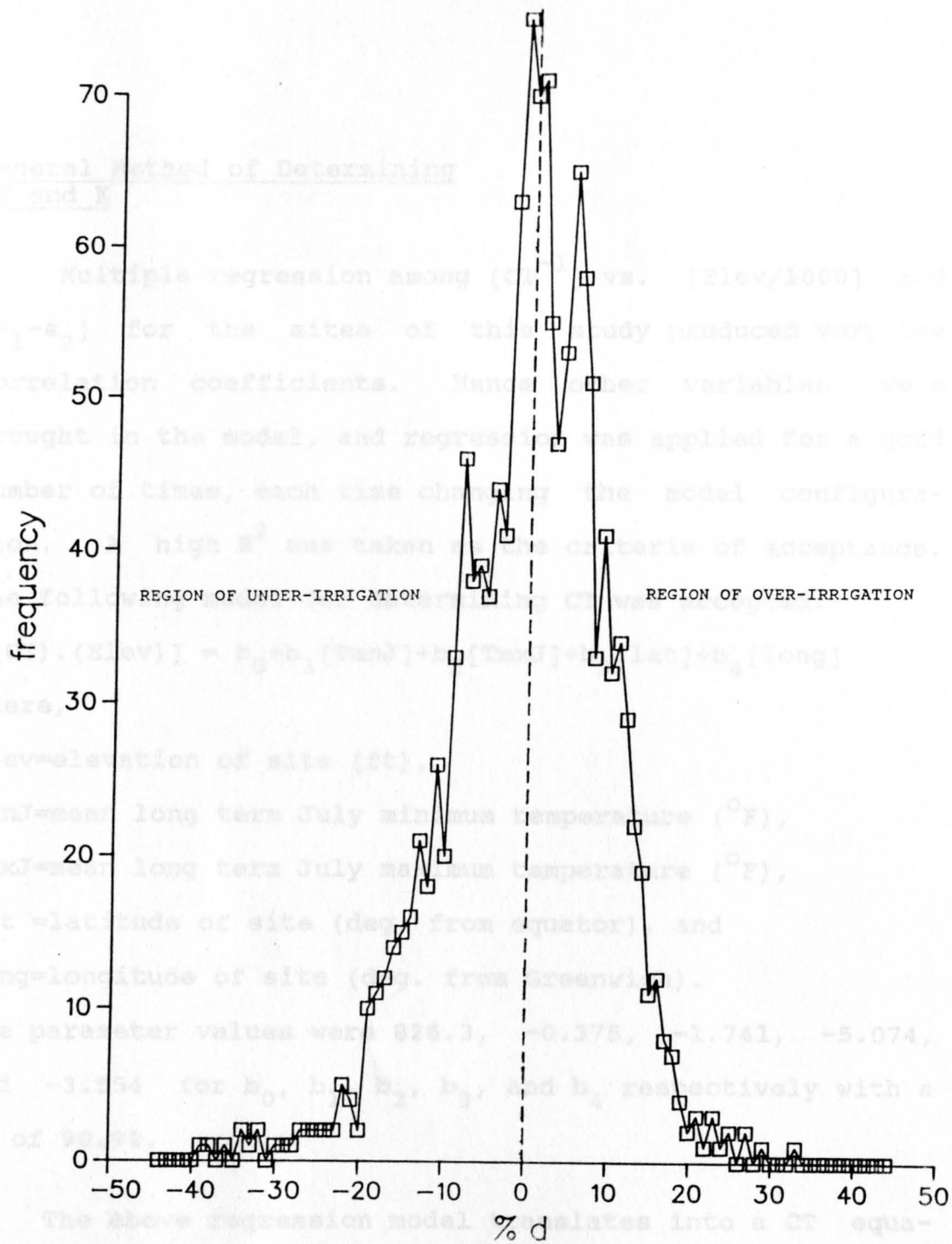


FIG. 12.- Plot of %d vs. Frequency of Occurrence for the Calibrated New Hargreaves ETr Equation as Compared to Penman ETr (All Sites Combined).

General Method of Determining  
CT and K

Multiple regression among  $[CT^{-1}]$  vs.  $[Elev/1000]$  and  $[e_1 - e_2]$  for the sites of this study produced very low correlation coefficients. Hence other variables were brought in the model, and regression was applied for a good number of times, each time changing the model configuration. A high  $R^2$  was taken as the criteria of acceptance. The following model for determining CT was accepted:

$$[(CT) \cdot (Elev)] = b_0 + b_1 [TmnJ] + b_2 [TmxJ] + b_3 [lat] + b_4 [long]$$

where,

Elev=elevation of site (ft),

TmnJ=mean long term July minimum temperature ( $^{\circ}F$ ),

TmxJ=mean long term July maximum temperature ( $^{\circ}F$ ),

Lat =latitude of site (deg. from equator), and

Long=longitude of site (deg. from Greenwich).

The parameter values were 826.3, -0.375, -1.741, -5.074, and -3.554 for  $b_0$ ,  $b_1$ ,  $b_2$ ,  $b_3$ , and  $b_4$  respectively with a  $R^2$  of 90.9%.

The above regression model translates into a CT equation as such:

$$CT = (b_0 + b_1 TmnJ + b_2 TmxJ + b_3 Lat + b_4 Long) / Elev \dots (20)$$

The same configuration was found to work best for the K of the New Hargreaves equation also. The respective  $b_0$ ,  $b_1$ ,  $b_2$ ,  $b_3$ , and  $b_4$  values were 85.2, 0.005, -0.228, -0.508, and -0.353 with a  $R^2$  of 85.2%.

This study, in compliance to its main objective, calibrated a Temperature-Radiation evapotranspiration equation ( $ETr = CT \times T \times Ra$ ) for Utah and some neighboring states by finding appropriate coefficients for specific sites.

The Temperature-Radiation equation was calibrated against the Kimberly version of the Penman ETr equation. Meteorological datafiles were used for the local calibration of the equation. In addition to the above equation, the New Hargreaves equation ( $ETr = K \times T \times TD^{0.5} \times Ra$ ) was similarly calibrated against the Kimberly version of the Penman ETr equation with the meteorological files of this study.

The accuracy of the above two equations can be understood by the fact that the Temperature-Radiation equation produced 5-day ETr estimates with 90.9% of the data lying within  $\pm 15\%$  of the Penman method, and the calibrated new Hargreaves equation produced results with 89.3% of data lying within  $\pm 15\%$  of the Penman method. These calibrated equations were meant for use by the Datapod field computer.

## SUMMARY AND CONCLUSIONS

This study, in compliance to its main objective, calibrated a Temperature-Radiation evapotranspiration equation ( $E_{Tr} = CT \times T \times R_s$ ) for Utah and some neighboring states by finding appropriate coefficients for specific sites.

The Temperature-Radiation equation was calibrated against the Kimberly version of the Penman  $E_{Tr}$  equation. Meteorological datafiles were used for the local calibration of the equation. In addition to the above equation, the New Hargreaves equation ( $E_{Tr} = K \times T \times TD^{0.5} \times R_a$ ) was similarly calibrated against the Kimberly version of the Penman  $E_{Tr}$  equation with the meteorological files of this study.

The accuracy of the above two equations can be understood by the fact that the Temperature-Radiation equation produced 5-day  $E_{Tr}$  estimates with 90.9% of the data lying within  $\pm 15\%$  of the Penman method, and the calibrated new Hargreaves equation produced results with 89.3% of data lying within  $\pm 15\%$  of the Penman method. These calibrated equations were meant for use by the Datapod field computer.



By considering the meteorological and site factors of the study sites, general methods for the estimation of the CT and K coefficients for any site in Utah and surroundings were developed. The necessary inputs were longterm July mean maximum and minimum temperatures, longitude, latitude, and elevation of sites. Appendix IV lists such inputs and the corresponding estimated CT and K coefficients, found by using equation (20), for various sites in the state of Utah.

Picking 29 farming communities in Utah, one corresponding to each county, a Datapod could be installed at representative sites for the purpose of estimating ETr. If the Temperature-Radiation ETr equation is utilized, a two channel Datapod, model DP219, would be required for each site. The cost of this particular model which comes with a TP10v temperature probe and a Licor pyranometer totals \$1,085 or \$31,465 as an initial cost for 29 sites. However if the New Hargreaves equation be substituted, a one channel Datapod model DP112 can be used instead. The DP112 comes equipped with a TP10v temperature probe and costs \$664 or \$19,256 for instrumentation of all the 29 sites; that is a savings of \$12,209 or \$421 per site. with consideration that the New Hargreaves equation produced ETr estimates within only 1.6% of the Temperature-Radiation meth-

od, usage of the DP112 Datapod model is recommended. However, since weather stations exist throughout the state of Utah with the recording of daily maximum and minimum temperatures, installation and purchase of the Datapod model DP112 may not even be required.

This will require installation of automatic lysimeters for use of neutron probe soil moisture maintenance and daily recording of data:

2) Crop coefficient curves for specific crops dominant in Utah can be developed.

3) The regression equation for estimation of  $R_0$  as a function of  $T_0$  and  $R_n$  needed as an input for the Row Hargreaves equation, can be further researched so as to produce even more accurate ETR estimates.

4) ETR vs.  $R_0$  curves for specific sites in Utah for prediction and adjustment of ETR forecasts can be developed.

5) Strategy and procedures can be thought of to gather weather information from dispersed areas to a central location in Utah for analysis. There may be possibility of using a satellite to monitor weathered stations throughout the state.

6) An implementation study can be done for utilizing the ETR equations calibrated in this study for developing

## RECOMMENDED FUTURE STUDIES

- 1) Calibration of the Penman ETr equation can be undertaken for sites in Utah against reliable measured data. This will require installation of sensitive lysimeters or use of neutron probes with careful maintenance and daily recording of data.
- 2) Crop coefficient curves for specific crops dominant in Utah can be developed.
- 3) The empirical equation for estimation of  $R_s$  as a function of TD and  $R_a$ , needed as an input for the New Hargreaves equation, can be further researched so as to produce even more accurate ETr estimates.
- 4) ETr vs. Time curves for specific sites in Utah for prediction and adjustment of ETr forecasts can be developed.
- 5) Strategy and methodology can be thought of to gather weather information from isolated areas to a central location in Utah for analysis. There may be possibility of using a satellite to monitor scattered stations throughout the state.
- 6) An implementation study can be done for utilizing the ETr equations calibrated in this study for developing

countries which have a potentially large irrigated agricultural sector and require accurate ETr estimations to assist in achieving high crop yields and in helping with water conservation.

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## APPENDICES

Appendix 1: Pennon EIR  
Question Elaboration

A brief description of equation 1 for estimating the inputs required for the Kimberly version of the Pennon EIR equation.

**APPENDIXES**



The Kimberly version of the Penman ETr equation is as follows:

$$ETr = \left( \frac{\Delta}{\Delta + \gamma} \right) (R_n + G) + \left( \frac{\gamma}{\Delta + \gamma} \right) 15.36 (W_e + W_e^2) (e_a - e_s) / CF$$

The following formulae can be used for the individual

### Appendix I: Penman ETr Equation Elaboration

A brief description of equations used for estimating the inputs required for the Kimberly version of the Penman ETr equation.

$$\Delta = 2.00(0.0041(T_f) + 0.676)^7 - 0.00116, \text{ for } T_f > -9.4^\circ\text{F}$$

$T_f$  is mean daily temperature in  $^\circ\text{F}$ .

$$\gamma = C_p P / (\lambda (M_w/M_a))$$

where  $C_p$  = Specific heat of air at constant pressure; it is taken as 0.240 Cal/gm/deg $^\circ\text{C}$ ,

$P = 1013 - 0.03216(\text{ELEV})$ ;  $P$  is air pressure in units of mb, and ELEV is site elevation in feet,

$\lambda$  is the latent heat of water estimated by:

$$\lambda = 595.5 - 0.305(T - 32), \text{ for } T \text{ in } ^\circ\text{F}$$

$M_w$  = Molecular weight of water equivalent to 18.016 gm/mole,

$M_a$  = Molecular weight of air equivalent to 28.966 gm/mole,  $(M_w/M_a) = 0.622$ .

$$R_n = (1 - \alpha) R_a - R_b$$

where  $\alpha$  is crop albedo.  $\alpha$  can be taken as 0.23 on the average for a green, actively growing crop at full height.

The Kimberly version of the Penman ETr equation stands as below:

$$ETr = [(\Delta / (\Delta + \gamma))(Rn + G) + (\gamma / (\Delta + \gamma))15.36(W1 + W2.U_2)(e_s - e_a)]CF$$

The following formulas can be used for the individual terms (17):

$$\Delta = 2.00(0.00738(Tc) + 0.8072)^7 - 0.00116, \text{ for } Tc > -23^\circ\text{C}.$$

$\Delta$  is in mb/ $^\circ\text{C}$  and Tc is mean daily temperature in  $^\circ\text{C}$ ,  
or:

$$\Delta = 2.00(0.0041(Tf) + 0.676)^7 - 0.00116, \text{ for } Tf > -9.4^\circ\text{F}.$$

Tf is mean daily temperature in  $^\circ\text{F}$ .

$$\gamma = Cp P / (\lambda (Mw/Ma))$$

where Cp = Specific heat of air at constant pressure; it is taken as 0.240 Cal/gm/deg $^\circ\text{C}$ ,

$P = 1013 - 0.03216(\text{ELEV})$ ; P is air pressure in units of mb, and ELEV is site elevation in feet,

$\lambda$  is the latent heat of water estimated by:

$$\lambda = 595.9 - 0.305(T - 32), \text{ for } T \text{ in } ^\circ\text{F},$$

Mw = Molecular weight of water equivalent to 18.016 gm/mole,

Ma = Molecular weight of air equivalent to 28.966 gm/mole,  $(Mw/Ma) \sim 0.622$ .

$$Rn = (1 - \alpha)Rs - Rb,$$

where  $\alpha$  is crop albedo.  $\alpha$  can be taken as 0.23 on the average for a green, actively growing crop at full height,

Rb is back radiation determined by:

$R_b = R_{bo}(a(R_s/R_{so}) + b)$ ,  $R_{bo}$  is back radiation for cloudless days and is approximated by:

$$R_{bo} = (a_1 + b_1 e_a) 11.71 \times 10^{-8} (T_a^4 + T_b^4) / 2.$$

$a, b, a_1,$  and  $b_1$  are empirical constants. Their recommended values for several locations are shown in Table 12.

$R_{so}$  is the clear day solar radiation for a particular day and site.  $R_{so}$  can be obtained from standard meteorological tables. For this study, empirical polynomial curves were used which gave a good approximation of  $R_{so}$  for each site.

$T_a$  and  $T_b$  are the daily maximum and minimum air temperatures in  $^{\circ}K$ .

$$G = 5(T_{pr} - T)$$

where  $T_{pr}$  = mean air temperature in  $^{\circ}F$  for a previous time period, usually the previous 3 days.  $T$  = mean daily air temperature in  $^{\circ}F$ .

$W_1$  and  $W_2$  are empirical wind parameters. Some suggested values for  $W_1$  and  $W_2$  are shown in Table 13.

$U_2$  = wind movement at 2 meter height from the ground.  $U_2$  can be approximated when the anemometer is at height  $x$  by:

$$U_2 = U_x (2/x)^{0.2}.$$

TABLE 12.- The a, b, a1, and b1 Empirical Constants Used for Estimating E<sub>h</sub> for Several Locations. From

$e_s$  = Saturation vapor pressure taken as the mean of values obtained at the daily maximum and minimum air temperatures,

$e_a$  = Mean actual vapor pressure taken as the saturation vapor pressure at the daily average dew point temperature. Vapor pressure at any given air temperature can be empirically approximated by:

$$e_s = \exp[21.3574 - 5336.0/(T^{\circ}\text{C} + 273.10)] \quad \text{or,}$$

$$e_s = \exp[21.3575 - 9604.8/(T^{\circ}\text{F} + 459.58)].$$

Care must be taken in using the above formulas. Where possible, local calibration should be done to determine exact figures.

TABLE 13.- Empirical Wind Parameters W1 and W2 for Several Reference Crops.

CROP	W1	W2	LOCATION AND/OR SOURCE
Short Green Crop	1.00	0.0100	Panama (31)
Alfalfa	0.75	0.0185	Kimberly, Idaho
Clipped Grass	1.00	0.0161	FAO Irrigation & Drainage Paper #24

TABLE 12.- The a, b, a1, and b1 Empirical Constants Used for Estimating  $R_{bo}$  for Several Locations. From Hill et al. (16).

a	b	a1	b1	LOCATION AND/OR SOURCE
1.35	-0.35	0.35	-0.046	Davis, CA
1.22	-0.18	0.325	-0.044	Kimberly, Idaho
1.20	-0.20	0.39	-0.05	Arid regions
1.10	-0.10	0.39	-0.05	Semihumid regions
1.00	0.00	0.39	-0.05	Humid regions
1.35	-0.35	0.34	-0.044	FAO Irrigation & Drainage paper #24

TABLE 13.- Empirical Wind Parameters W1 and W2 for Several Reference Crops.

CROP	W1	W2	LOCATION AND/OR SOURCE
Short Green Crop	1.00	0.0100	Penman (31)
Alfalfa	0.75	0.0185	Kimberly, Idaho
Clipped Grass	1.00	0.0161	FAO Irrigation & Drainage Paper #24



\*\*\*\*\*  
 \* "CROPIN.FOR" \*  
 \*\*\*\*\*

THIS IS A FORTRAN PROGRAM NAMED "CROPIN.FOR". IT ESTIMATES  
 ALFALFA REFERENCE EVAPOTRANSPIRATION IN INCHES/DAY THROUGH  
 THE KIRKBYLY VERSION OF THE PENMAN ETc EQUATION. THE MODEL  
 REQUIRES INPUT DATA AND SITE SPECIFICATIONS AND METEOROLOGICAL  
 DATA WHICH THE PROGRAM READS FROM A GIVEN SITE DATAFILE.

PROGRAM WAS WRITTEN AND DEVELOPED WITHIN THE MECHANICAL  
 AND CIVIL ENGINEERING DEPARTMENT,  
 IOWA STATE UNIVERSITY,  
 UNDER THE DIRECTION OF: DR. F. W. HILZ

Appendix II: Major Programs and  
Sample Datafile Listings

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 VARIABLE AND CONSTANT IDENTIFICATION  
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1000	*	A, AA= VARIABLES REQUIRED FOR THE ESTIMATION OF NET RADIATION. AN INPUT OF THE PENMAN ETc EQUATION.
1050	*	AT, A1, AT, A2, A3, A4, A5, A6, A7= POLYNOMIAL VARIABLES USED TO ESTIMATE THE CLEAR DAY SOLAR RADIATION (RS00) OF THE SITE.
1100	*	ALAM= LATENT HEAT OF WATER, APPROXIMATELY 585 CAL/(CUBIC CM).
1200	*	B, B1= SIMILAR TO A AND AA.
1300	*	CONV= THE CONVERSION FACTOR REQUIRED TO CONVERT UNITS OF GIVEN THE MONTH AND DAY.
1400	*	LANGLEY IN/INCHES OF EQUIVALENT WATER EVAPORATION. CONV IS EQUAL TO 0.000875 IN/LANGLEYS FOR STANDARD CONDITIONS.
1500	*	DELTA= ESTIMATE OF THE SLOPE OF SATURATION VAPOR PRESSURE-TEMPERATURE CURVE AT THE AIR TEMPERATURE IN mb/deg C.
1600	*	ELEV= ELEVATION OF SITE IN FEET.
1700	*	E1, E2= SATURATION VAPOR PRESSURES FOR THE WARMEST MONTH OF MONTH, ASSUMED TO BE JULY, AT THE MEAN MAXIMUM AND MEAN MINIMUM TEMPERATURES OF THE SITE RESPECTIVELY.
1800	*	FMT= THE FORMAT BY WHICH THE METEOROLOGICAL DATA FILE ARE.
1900	*	IDAYS, IDAYS= BEGINNING AND ENDING JULIAN DAY OF THE DATA.
2000	*	GAMA= PSYCHROMETRIC CONSTANT IN mb/deg C.
2100	*	MON, END= BEGINNING AND ENDING MONTH OF THE NET. DATA.
2200	*	PENET(1)= DAILY CALCULATED PENMAN ETc IN INCHES/DAY.
2300	*	PENET(5)= FIVE-DAY SUM OF PENMAN ETc READ FROM THE DATA.
2400	*	P= AVERAGE AIR PRESSURE ESTIMATED BY THE AVERAGE TEMPERATURE IN mb/g OR mb.
2500	*	RS(I)= DAILY SOLAR RADIATION DATA IN LANGLEYS.
2600	*	STNG= STATION NAME.
2700	*	SVPAT= AVERAGE DAILY SATURATION VAPOR PRESSURE ESTIMATED BY THE AVERAGE DAILY TEMPERATURE.
2800	*	TMX(I)= DAILY MAXIMUM AIR TEMPERATURE IN DEG F IN DATA FILE.

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100 *****
200 *
300 *           " CRPSIM.FOR "
400 *
500 *****
600 *
700 *   THIS IS A FORTRAN PROGRAM NAMED "CRPSIM.FOR". IT ESTIMATES
800 *   ALFALFA REFERENCE EVAPOTRANSPIRATION IN INCHES/DAY THROUGH
900 *   THE KIMBERLY VERSION OF THE PENMAN ETr EQUATION. THE NECE-
1000 *   SSARY INPUT DATA ARE SITE SPECIFICATIONS AND METEOROLOGICAL
1100 *   DATA WHICH THE PROGRAM READS FROM A GIVEN SITE'S DATAFILE.
1200 *
1300 *****
1400 *
1500 *   PROGRAM WAS WRITTEN AND DEVELOPED WITHIN THE AGRICULTURAL
1600 *   AND IRRIGATION ENGINEERING DEPARTMENT,
1700 *   UTAH STATE UNIVERSITY,
1800 *   UNDER THE DIRECTION OF: DR R. W. HILL
1900 *   MODIFIED BY: P. FOROUGH
2000 *   DECEMBER 1984
2100 *
2200 *****
2300 *
2400 *           VARIABLE AND CONSTANT IDENTIFICATION
2500 *
2600 *   A,AA= VARIABLES REQUIRED FOR THE ESTIMATION OF NET RADIA-
2700 *   TION, AN INPUT OF THE PENMAN ETr EQUATION.
2800 *   AO,A1,A2,A3,A4,A5,A6,A7= POLINOMIAL VARIABLES USED TO ESTI-
2900 *   MATE THE CLEAR DAY SOLAR RADIATION (RS0) OF THE SITE.
3000 *
3100 *   ALAM= LATENT HEAT OF WATER, APPROXIMATELY 585 CAL/(CUBIC CM).
3200 *
3300 *   B,BB= SIMILAR TO A AND AA.
3400 *
3500 *   CONV= THE CONVERSION FACTOR REQUIRED TO CONVERT UNITS OF
3600 *   GIVEN THE MONTH AND DAY.
3700 *   LANGLEYS INTO INCHES OF EQUIVALENT WATER EVAPORATION, CONV
3800 *   IS EQUAL TO 0.000673 IN/LANGLEYS FOR STANDARD CONDITIONS.
3900 *
4000 *   DELTA= ESTIMATE OF THE SLOPE OF SATURATION VAPOR PRESSURE-
4100 *   TEMPERATURE CURVE AT THE AIR TEMPERATURE IN mb/deg C.
4200 *
4300 *   ELEV= ELEVATION OF SITE IN FEET.
4400 *
4500 *   E1,E2= SATURATION VAPOR PRESSURES FOR THE WARMEST MONTH OF
4600 *   MONTH, ASSUMED TO BE JULY, AT THE MEAN MAXIMUM AND MEAN
4700 *   MINIMUM TEMPERATURES OF THE SITE RESPECTIVELY.
4800 *
4900 *   FMT= THE FORMAT AT WHICH THE METEOROLOGICAL DATA FILE ARE.
5000 *
5100 *   IDAYB,IDAYE= BEGINNING AND ENDING JULIAN DAY OF THE DATA.
5200 *
5300 *   GAMA= PSYCHROMETRIC CONSTANT IN mb/deg C.
5400 *
5500 *   MNB,MNE= BEGINNING AND ENDING MONTH OF THE MET. DATA.
5600 *
5700 *   PENET(I)= DAILY CALCULATED PENMAN ETr IN INCHES/DAY.
5800 *   PENET5(I)= FIVE-DAY SUM OF PENMAN ETr READ FROM THE DATA.
5900 *
6000 *   P= AVERAGE AIR PRESSURE ESTIMATED BY THE AVERAGE TEMPERA-
6100 *   TURE IN mmHg OR mb.
6200 *
6300 *   RS(I)= DAILY SOLAR RADITION DATA IN LANGLEYS.
6400 *
6500 *   STNM= STATION NAME.
6600 *
6700 *   SVPVAV= AVERAGE DAILY SATURATION VAPOR PRESSURE ESTIMATED BY
6800 *   THE AVERAGE DAILY TEMPERATURE.
6900 *
7000 *   TMX(I)= DAILY MAXIMUM AIR TEMPERATURE IN DEG F IN DATA FILE.

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7100 *      TMN(I)= DAILY MINIMUM AIR TEMPERATURE. *
7200 *      TWB(I)= DAILY WET BULB TEMPERATURE. *
7300 *      TDB(I)= DAILY DRY BULB TEMPERATURE. *
7400 *      TAV(I)= DAILY AVERAGE AIR TEMPERATURE. *
7500 *      TMNW, TMXW= AVERAGE LONG TERM MINIMUM AND MAXIMUM JULY TEM- *
7600 *      PERATURE RESPECTIVELY. *
7700 *      TEMCN= TEMPERATURE ADJUSTMENT CONSTANT. *
7800 * *
7900 *      W1, W2= WIND PARAMETERS USED IN THE PENMAN EQUATION. *
8000 *      WHT= ANEMOMETER HEIGHT IN METERS. *
8100 *      WIND(I)= DAILY WIND RUN IN MILES. *
8200 *      WINDCN= WIND READING ADJUSTMENT CONSTANT. *
8300 *      WINDLIM= WIND LIMIT SET AT 100 MILES/DAY FOR THIS STUDY. *
8400 * *
8500 *      XNDP= NUMBER OF DAY OF DATA. *
8600 * *
8700 *****
8800 *
8900      DIMENSION TMX(250), TMN(250), TWB(250), TDB(250), PPT(250),
9000 +WIND(250), RS(250), TAV(250), PENET(250), CONV(250)
9100      CHARACTER STNM*25, FMT*50
9200      OPEN(5, FILE='PENINP', STATUS='OLD')
9300      OPEN(8, FILE='PEN', STATUS='NEW')
9400 *
9500 ***** SITE SPECIFICATIONS ARE READ.
9600 *
9700      READ(5, 10) STNM, ELEV
9800      10  FORMAT(/A25, 1X, F5.0/////)
9900      READ(5, 20) TMXW, TMNW, A, B, AA, BB, W1, W2, WHT, WNDLIM
10000     20  FORMAT(2F10.0/3X, 12F6.0)
10100     READ(5, 30) A0, A1, A2, A3, A4, A5, A6
10200     30  FORMAT(5X, 7E10.0)
10300     READ(5, 40) MNB, IDAYB, MNE, IDAYE
10400     40  FORMAT(4I5)
10500     READ(5, 45) RSCN, WINDCN, PPCN, PANCN, TEMCN
10600     IF(WINDCN.EQ.0) WINDCN=1
10700     45  FORMAT(5F10.0)
10800     READ(5, 50) FMT
10900     WRITE(99, 51) TMXW, TMNW, A, B, AA, BB, W1, W2, WHT, WNDLIM,
11000 +A0, A1, A2, A3, A4, A5
11100     51  FORMAT(1X'TXW='F5.2' TNW='F5.2' A='F5.3' B='F5.3' AA='F5.3
11200 + ' BB='F5.3/3X'W1='F5.3' W2='F5.3' WHT='F4.2' WNDLIM='F5.1/
11300 +3X'A0-A6='7(E10.3, 2X))
11400     50  FORMAT(A50)
11500     CALL JULDAY(MNB, IDAYB, AJDB)
11600     CALL JULDAY(MNE, IDAYE, AJDE)
11700     XNDP=AJDE - AJDB + 1.
11800     TMXW=(TMXW-32.)*5./9.
11900     TMNW=(TMNW-32.)*5./9.
12000     CALL SAVAPR(TMXX, SVPTXW)
12100     CALL SAVAPR(TMNW, SVPTNW)
12200     SVPDIF=SVPTXW - SVPTNW
12300     TAVW=(TMXW + TMNW)/2.
12400     CP=.240
12500     P=1013. - 0.03217*ELEV
12600     ALAM1=595.9 - .549*TAVW
12700     CONV1=1/(2.54*ALAM1)
12800     AJD=AJDB - 1.
12900 *
13000 ***** DAILY METEOROLOGICAL DATA ARE READ.
13100 *
13200     DO 80 I=1, XNDP
13300     READ(5, FMT) TMX(I), TMN(I), TWB(I), TDB(I), PPT(I), WIND(I), RS(I)
13400     TAV(I)=(TMN(I) + TMX(I))/2
13500     IF(TEMCN.EQ.1) GOTO 55
13600     TMX(I)=(TMX(I) - 32.)*5./9.
13700     TMN(I)=(TMN(I) - 32.)*5./9.
13800     IF(TDB(I).EQ.0) GOTO 52
13900     TDB(I)=(TDB(I) - 32.)*5./9.
14000     52  TWB(I)=(TWB(I) - 32.)*5./9.

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14100      TAV(I)=(TAV(I)-32.)*5./9.
14200
14300      55  DEL=0.00738*TAV(I) + 0.8072
14400      DEL7=DEL*DEL*DEL*DEL*DEL*DEL*DEL
14500      DELTA=2.00* DEL7 - .00116
14600      ALAM=595.9 - .549*TAV(I)
14700      GAMA=CP*P/(.622*ALAM)
14800      DELGAM=DELTA/(DELTA + GAMA)
14900      CONV(I)= 1/(2.54*ALAM)
15000
15100      CALL SAVAPR(TWB(I),SVPWB)
15200      SVPDP = SVPWB - GAMA*(TDB(I)-TWB(I))
15300      IF(TDB(I).EQ.0)SVPDP=SVPWB
15400      CALL SAVAPR(TMX(I),SVPTMX)
15500      CALL SAVAPR(TMN(I),SVPTMN)
15600      SPPAV=(SVPTMX + SVPTMN)/2.
15700
15800      TXK=TMX(I) + 273.1
15900      TNK=TMN(I) + 273.1
16000      TXK4=TXK*TXK*TXK*TXK
16100      TNK4=TNK*TNK*TNK*TNK
16200      TAVK4=(TXK4+TNK4)/2.
16300      RBO = (AA + BB*SQRT(SVPDP))*11.71E-8*TAVK4
16400      AJD=AJD+1
16500      RSO=A0+AJD*(A1+AJD*(A2+AJD*(A3+AJD*(A4+AJD*(A5+AJD*A6))))
16600      RB = RBO*(B + A*RS(I)/RSO)
16700      RN = 0.77*RS(I) - RB
16800
16900      IF(TAV(I-1).EQ.0)TAV(I-1)=TAV(I)
17000      IF(TAV(I-2).EQ.0)TAV(I-2)=TAV(I-1)
17100      IF(TAV(I-3).EQ.0)TAV(I-3)=TAV(I-2)
17200      TPR=(TAV(I-3) + TAV(I-2) + TAV(I-1))/3.
17300      TPR=32.+TPR*9./5.
17400      G=(TPR - 32.-TAV(I)*9./5.)*5.
17500
17600      WIND(I)=WIND(I)*WINDCN
17700      IF(WIND(I).GT.100.)WIND(I)=100.
17800      IF(WHT.EQ.2)THEN
17900          U2=WIND(I)
18000      ELSE
18100          U2=WIND(I)*(2./WHT)**.2
18200      ENDIF
18300      *
18400      **** DAILY PENMAN ETr IS CALCULATED.
18500      *
18600      PENET(I)= (DELGAM*(RN+G) +
18700      + (1-DELGAM)*15.36*(W1 +W2*U2)*(SPPAV-SVPDP) ) *CONV(I)
18800      SUMPEN=SUMPEN + PENET(I)
18900      CALL ROOZ(AJD,IMN,IDAY)
19000      80  CONTINUE
19100      *
19200      **** PROGRAM RESULTS ARE PRINTED.
19300      *
19400      WRITE(8,200)STNM,ELEV,SUMPEN
19500      200  FORMAT(5X,A25' ELEVATION=' F6.1/5X'SEASONAL PENMAN ET= 'F5.2/)
19600      WRITE(8,250)CTOBJ2, TXOBJ2, FIX, PFIX, SUBOBJ1, SEASOBJ2
19700      250  FORMAT(/5X'FOR LOWEST OBJ2: CT= 'F6.4' AND TX= 'F5.1/
19800      +5X'OBJ2= 'F7.4' (OBJ2/SUMPEN)*100= 'F5.2'%/5X'OBJ1= 'F7.4/
19900      +5X'SUMJH= 'F5.2/)
20000      WRITE(8,300)CTOBJ1, TXOBJ1, HOLD, SEASOBJ1
20100      300  FORMAT(5X'FOR LOWEST OBJ1: CT= 'F6.4' AND TX= 'F5.1/
20200      +5X'OBJ1= 'F7.4/5X'SUMJH= 'F5.2/5X'SUM PENMAN  PENMANET')
20300      DO I=1,(XNDP+2),5
20400          PENET5=PENET(I)+PENET(I+1)+PENET(I+2)+PENET(I+3)+PENET(I+4)
20500          SPENET5=SPENET5+PENET5
20600          WRITE(8,500)SPENET5,PENET5
20700      ENDDO
20800      WRITE(8,400)CTZOBJ2, ZOBJ2, PZOBJ2, SEASZOBJ2
20900      400  FORMAT(5X'FOR TX OF ZERO: CT= 'F6.4/5X'OBJ2= 'F7.4
21000      +' (OBJ2/SUMPEN)*100= 'F5.2'%/5X'SUMJH= 'F5.2/

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21100 .....+5X'SUM JHET  JHET')
21200 500  FORMAT(7X,F5.2,8X,F5.2)
21300      STOP
21400      END
21500 .....
21600 *
21700 ****  SUBROUTINE SAVAPR ESTIMATES THE SATURATION VAPOR PRESSURE
21800 *      FOR A GIVEN TEMPERATURE.
21900 *
22000      SUBROUTINE SAVAPR(TC,SVP)
22100      SVP= 1.3329*EXP(21.07 - 5336./(TC+273.1))
22200      RETURN
22300      END
22400 *
22500 ****  SUBROUTINE JULDAY CALCULATES THE JULIAN DAY OF THE YEAR
22600 *
22700      SUBROUTINE JULDAY(IMN,IDAY,AJD)
22800      DIMENSION MN(12)
22900      DATA MN/31,28,31,30,31,30,31,31,30,31,30,31/
23000      SUM=0
23100      DO 20 I=1,IMN
23200      IF (IMN.EQ.I)GOTO 10
23300      SUM=SUM+MN(I)
23400 10  AJD=SUM+IDAY
23500 20  CONTINUE
23600      RETURN
23700      END
23800 *
23900 ****  SUBROUTINE ROOZ CALCULATES THE DAY AND MONTH GIVEN THE
24000 *      JULIAN DAY.
24100 *
24200      SUBROUTINE ROOZ(AJD,IMN,IDAY)
24300      DIMENSION MN(12)
24400      DATA MN/31,28,31,30,31,30,31,31,30,31,30,31/
24500      SUM=0
24600      DO 10 J=1,12
24700      SUM=SUM+MN(J)
24800      IF ((AJD-SUM).LE.0)GOTO 20
24900 10  CONTINUE
25000 20  IMN=J
25100      IDAY=AJD - SUM + MN(J)
25200      RETURN
25300      END
25400 *
25500 *      END OF PROGRAM
25600 *
25700 *****

```

```

100 *****
200 *
300 *           " TEMPRAD.FOR "
400 *
500 *****
600 *
700 *   THIS IS A FORTRAN PROGRAM NAMED "TEMPRAD.FOR". IT READS THE
800 *   METEOROLOGICAL DATA OF A SITE AND THE PENMAN ETr VALUES
900 *   GENERATED FOR THAT SITE FROM A DATAFILE. TEMPRAD THEN AT-
1000 *   TEMPTS TO CALIBRATE THE TEMPERATURE-RADIATION ETr EQUATION
1100 *   FOR THAT SITE.
1200 *
1300 *   WRITTEN BY: PAYAM FOROUGHI
1400 *   UTAH STATE UNIVERSITY
1500 *   DECEMBER 1984
1600 *
1700 *****
1800 *
1900   REAL JHET(365), JHET5
2000   DIMENSION TMX(365), TMN(365), TWB(365), TDB(365), PPT(365), TAVF(365),
2100   +WIND(365), RS(365), TAV(365), PENET(365), CONV(365), ETJH(365),
2200   +PENET5(365), EVAP(365)
2300   CHARACTER STNM*25, FMT*70
2400   OPEN(5, FILE='PENINP', STATUS='OLD')
2500   OPEN(8, FILE='JH', STATUS='NEW')
2600 *
2700 ****   SITE SPECIFICATIONS ARE READ.
2800 *
2900   READ(5, 10) STNM, ELEV
3000   10   FORMAT(/A25, 1X, F5.0////)
3100   READ(5, 15) TMXW, TMNW, A, B, AA, BB, W1, W2, WHT, WINDLIM
3200   15   FORMAT(2F10.0/3X, 12F6.0)
3300   READ(5, 20) K, A0, A1, A2, A3, A4, A5, A6
3400   20   FORMAT(3X, I2, 7E10.0)
3500   READ(5, 30) MNB, IDAYB, MNE, IDAYE
3600   30   FORMAT(4I5)
3700   READ(5, 40) RSCN, WINDCN, PPCN, FANCN, TEMCN
3800   IF(RSCN.LE.0) RSCN=1.00
3900   IF(WINDCN.LE..001) WINDCN=1.00
4000   40   FORMAT(5F10.0)
4100   READ(5, 45) FMT
4200   45   FORMAT(A70)
4300   CALL JULDAY(MNB, IDAYB, JDB)
4400   CALL JULDAY(MNE, IDAYE, JDE)
4500   XNDP=FLOAT(JDE-JDB+1)
4600 *
4700 ****   DAILY METEOROLOGICAL DATA ARE READ.
4800 *
4900   DO 50 I=JDB, JDE
5000   READ(5, FMT) TMX(I), TMN(I), TWB(I), TDB(I), PPT(I), WIND(I), RS(I)
5100   TAV(I)=(TMX(I)+TMN(I))*0.5
5200   RS(I)=RS(I)*RSCN
5300   IF(TEMCN.EQ.1.) THEN
5400     TAV(I)=32. + TAV(I)*9./5.
5500     TMX(I)=32. + TMX(I)*9./5.
5600     TMN(I)=32. + TMN(I)*9./5.
5700   ENDIF
5800   SPPT=SPPT+PPT(I)
5900   STMN=STMN+TMN(I)
6000   STMX=STMX+TMX(I)
6100   CALL SAVAPR(TMN(I), SVPTMX)
6200   CALL SAVAPR(TMN(I), SVPTMN)
6300   SVPVAV=(SVPTMX+SVPTMN)/2.
6400   SSVPAV=SSVPAV+SVPVAV
6500   IF((I.GE.182).AND.(I.LE.212)) THEN
6600     SJULSVP=SJULSVP+SVPVAV
6700     SJULTMN=SJULTMN+TMN(I)
6800     SJULTMX=SJULTMX+TMX(I)
6900   ENDIF
7000   WIND(I)=WIND(I)*WINDCN

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7100      IF(WHT.EQ.2.) THEN
7200          U2=WIND(I)
7300          ELSE
7400              U2=WIND(I)*(2./WHT)**.2
7500      ENDIF
7600      IF((I.GT.121).AND.(I.LE.151)) SWMAY=SWMAY+U2
7700      IF((I.GT.151).AND.(I.LE.181)) SWJUN=SWJUN+U2
7800      IF((I.GT.181).AND.(I.LE.212)) SWJUL=SWJUL+U2
7900      IF((I.GT.212).AND.(I.LE.243)) SWAUG=SWAUG+U2
8000      IF((I.GT.243).AND.(I.LE.273)) SWSEP=SWSEP+U2
8100  50      CONTINUE
8200      READ(5,55) SUMPEN
8300  55      FORMAT(/5X,F5.2)
8400      DO 60 I=JDB,JDE,5
8500  *
8600  ****   PENMAN 5-DAY SUM ETr VALUES ARE READ.
8700  *
8800  60      READ(5,65) PENET5(I)
8900  65      FORMAT(5X,F5.2)
9000      WRITE(8,68) STNM,MNB, IDAYB,MNE, IDAYE, ELEV, SUMPEN
9100  68      FORMAT(5X,A25,2X,I2','I2'-'I2'/'I2
9200      +//5X'ELEVATION='F5.0/5X'SEASONAL PENMAN ET='F5.2/)
9300      XNDP=JDE - JDB + 1.
9400  *
9500  ****   JENSEN-HAISE 1970 ET EQUATION CONSTANTS ARE CALCULATED.
9600  *
9700      C1=68. - (3.6*ELEV/1000.)
9800      C2= 13.
9900      CALL SAVAPR(TMXW,E2)
10000     CALL SAVAPR(TMNW,E1)
10100     SVPDIF=E2-E1
10200     CH=50./SVPDIF
10300     CT70= 1./(C1 + C2*CH)
10400     TX70= 27.5 - .25*SVPDIF - ELEV/1000.
10500
10600     TAVW=(TMXW + TMNW)/2.
10700     P=1013. - 0.03217*ELEV
10800     WRITE(91,*)STNM,' PENMAN ET'
10900     WRITE(92,*)STNM,' JEN-HAS 1970 ET'
11000     WRITE(93,*)STNM,' TEM-RAD (DATAPOD) ET'
11100
11200  *
11300  ****   CALIBRATION OF THE TEMPERATURE-RADIATION ETr EQUATION IS BEGUN.
11400  *
11500     HOLD=1000.
11600     FIX=1000.
11700     ZOBJ2=1000.
11800     PRINT*,'SUMPEN=',SUMPEN
11900     DO 100 CT=.0005,.0200,.0005
12000     DO 100 TX=-30.,20.,.5
12100     SUMJH=0.0
12200     OBJ1=0.0
12300     JHET5=0.0
12400     DO 90 I=JDB,JDE
12500     EVAP(I)=CT*(TAV(I)-TX)*RS(I)*.000673
12600     SUMJH=SUMJH+EVAP(I)
12700  90     CONTINUE
12800     DO 95 I=JDB,JDE,5
12900     JHET5=EVAP(I)+ EVAP(I+1)+ EVAP(I+2)+ EVAP(I+3)+ EVAP(I+4)
13000     DIFF= ABS(PENET5(I) - JHET5)
13100     OBJ1= OBJ1 + DIFF
13200  95     CONTINUE
13300     SEADIF= ABS(SUMPEN - SUMJH)
13400     OBJ2 = OBJ1 + SEADIF
13500     POBJ2=(OBJ2/SUMPEN)*100.
13600     IF(OBJ1.LT.HOLD) THEN
13700         HOLD=OBJ1
13800         CTOBJ1=CT
13900         TXOBJ1=TX
14000         SEASOBJ1=SUMJH

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14100      ENDIF
14200      IF (OBJ2.LT.FIX) THEN
14300          FIX=OBJ2
14400          PFIX=POBJ2
14500          SUBOBJ1=OBJ1
14600          CTOBJ2=CT
14700          TXOBJ2=TX
14800          SEASOBJ2=SUMJH
14900      ENDIF
15000      IF ((TX.EQ.0).AND.(OBJ2.LT.ZOBJ2)) THEN
15100          ZOBJ1=OBJ1
15200          ZOBJ2=OBJ2
15300          PZOBJ2=POBJ2
15400          CTZOBJ2=CT
15500          SEASZOBJ2=SUMJH
15600      ENDIF
15700  100    CONTINUE
15800
15900      PRINT*, 'CT=', CTZOBJ2, ' WHEN TX=0 WITH BEST OBJ'
16000      PRINT*, 'ENTER THE CT AND TX TO BE USED IN THE DATAPOD:'
16100      READ*, CTB, TXB
16200      JIM=0
16300          CT=CT70
16400          TX=TX70
16500      IF (JIM.EQ.0) GOTO 120
16600  115    CT=CTB
16700          TX=TXB
16800  120    DO 125 I=JDB, JDE
16900  125    ETJH(I)=CT*(TAV(I) - TX)*RS(I)*.000673
17000      DO 130 I=JDB, JDE, 5
17100          JHET5=ETJH(I)+ETJH(I+1)+ETJH(I+2)+ETJH(I+3)+ETJH(I+4)
17200          IF (JIM.EQ.0) THEN
17300              OBJ170=OBJ170+ABS(PENET5(I)-JHET5)
17400              SJHET70=SJHET70+JHET5
17500              SPENET5=SPENET5+PENET5(I)
17600              WRITE(91,300) I, SPENET5, PENET5(I)
17700              WRITE(92,300) I, SJHET70, JHET5
17800          ENDIF
17900          IF (JIM.EQ.1) THEN
18000              OBJ1BST=OBJ1BST+ABS(PENET5(I)-JHET5)
18100              SJHETZ=SJHETZ+JHET5
18200              WRITE(93,300) I, SJHETZ, JHET5
18300          ENDIF
18400  130    CONTINUE
18500      IF (JIM.EQ.0) THEN
18600          DIFF70=ABS(SUMPEN-SJHET70)
18700          OBJ270=OBJ170+DIFF70
18800          POBJ270=(OBJ270/SUMPEN)*100.
18900      ENDIF
19000      IF (JIM.EQ.1) THEN
19100          DIFFBST=ABS(SUMPEN-SJHETZ)
19200          OBJ2BST=OBJ1BST+DIFFBST
19300          POBJ2BST=(OBJ2BST/SUMPEN)*100.
19400          GOTO 175
19500      ENDIF
19600      IF (CTB.EQ.9) THEN
19700          CONTINUE
19800      ELSE
19900          JIM=1
20000          GOTO 115
20100      ENDIF
20200  175    TMXWC=(TMXW-32.)*5./9.
20300          TMNWC=(TMNW-32.)*5./9.
20400      *
20500      **** PROGRAM OUTPUT IS PRINTED.
20600      *
20700          WRITE(8,250) CTOBJ2, TXOBJ2, FIX, PFIX, SUBOBJ1, SEASOBJ2
20800  250    FORMAT(5X'FOR LOWEST OBJ2: CT= 'F7.5' AND TX= 'F5.1/
20900          +5X'OBJ2= 'F7.4', (OBJ2/SUMPEN)= 'F5.2'%/5X'OBJ1= 'F7.4/
21000          +5X'SUMJH= 'F5.2/)

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21100 300  FORMAT(1X,I3,1X,F5.2,1X,F5.2)
21200      WRITE(8,350)CTOBJ1,TXOBJ1,HOLD,SEASOBJ1
21300 350  FORMAT(5X'FOR LOWEST OBJ1: CT= 'F7.5' AND TX= 'F5.1/
21400      +5X'OBJ1= 'F7.4/5X'SUMJH= 'F5.2)
21500      WRITE(8,400)CTZOBJ2,ZOBJ1,ZOBJ2,PZOBJ2,SEASZOBJ2
21600 400  FORMAT(1X,55('*')/1X*'53X'*/1X
21700      +'*3X'FOR TX OF ZERO: CT= 'F7.5,T56'*/1X*'3X'OBJ1='F7.4,T56'*'
21800      +/1X*'3X'OBJ2= 'F7.4', (OBJ2/SUMPEN)= ',F6.2'*/1X
21900      +'*3X'SUMJH= 'F5.2,T56'*/1X*'53X'*/1X,55('*')/)
22000      IF(CTB.EQ.9)GOTO 575
22100      WRITE(8,550)CTB,TXB,OBJ1BST,OBJ2BST,POBJ2BST,SJHETZ
22200 550  FORMAT(5X'DATA POD WILL USE: CT = 'F7.5' AND TX = 'F5.1/5X'OBJ1=
22300      + 'F7.4/5X'OBJ2= 'F7.4', (OBJ2/SUMPEN)= 'F6.2'*/5X'SUMJH= 'F5.2//)
22400 575  WRITE(8,500)CT70,TX70,OBJ170,OBJ270,POBJ270,SJHET70
22500 500  FORMAT(5X'FOR J-H 1970: CT= 'F7.5' AND TX= 'F5.2/5X'OBJ1= 'F7.4/
22600      +5X'OBJ2= 'F7.4', (OBJ2/SUMPEN)= 'F6.2'*/5X'SUMJH= 'F5.2//)
22700      TMNAV=STMN/XNDP
22800      TMNAVC=(TMNAV - 32.)*5./9.
22900      TMXAV=STMX/XNDP
23000      TMXAVC=(TMXAV-32.)*5./9.
23100      TDF=TMXAV-TMNAV
23200      TDFC=TMXAVC-TMNAVC
23300      TAVE=(TMNAV+TMXAV)*.5
23400      TAVEC=(TMNAVC+TMXAVC)*.5
23500      AJULTMN=SJULTMN/31.
23600      AJULTMNC=(AJULTMN-32.)*5./9.
23700      AJULTMX=SJULTMX/31.
23800      AJULTMXC=(AJULTMX-32.)*5./9.
23900      AJULTAV=(AJULTMN+AJULTMX)*.5
24000      AJULTAVC=(AJULTMNC+AJULTMNC)*.5
24100      AJULTDFC=AJULTMXC - AJULTMNC
24200      SEASSVDP=SSVPAV/XNDP
24300      AJULSVPD=SJULSVP/31.
24400      AJULTDF=AJULTMX-AJULTMN
24500      WRITE(8,600)TMNAV,TMNAVC,AJULTMN,AJULTMNC,TMXAV,TMXAVC,AJULTMX,
24600      +AJULTMXC,TDF,TDFC,AJULTDF,AJULTDFC,TAVE,TAVEC,AJULTAV,AJULTAVC,
24700      +SPPT,SEASSVDP,AJULSVPD,TMXW,TMNW,(TMXW-TMNW),TMXWC,TMNWC,
24800      +(TMXWC-TMNWC),E2,E1,(E2-E1)
24900 600  FORMAT(24X'S U M M A R Y'/18X'SEASONAL'13X'JULY'/5X
25000      +'AVE. TMN='F6.2' F, 'F6.2' C 'F6.2' C, 'F6.2' C'/5X
25100      +'AVE. TMX='F6.2' F, 'F6.2' C 'F6.2' F, 'F6.2' C'/5X
25200      +' DIFF.='F6.2' F, 'F6.2' C 'F6.2' F, 'F6.2' C'/5X
25300      +' TAV = 'F6.2' F, 'F6.2' C 'F6.2' F, 'F6.2' C'/5X
25400      +' PRECIP='F5.2' INCHES'/
25500      +'      E2-E1 = 'F7.4' mb      E2-E1 = 'F7.4' mb//
25600      +'      MEAN HISTORICAL JULY VALUES'/
25700      +'      TMX: 'F6.2' - TMN: 'F6.2' = 'F6.2' C'/
25800      +'      TMX: 'F6.2' - TMN: 'F6.2' = 'F6.2' F'/
25900      +'      E2: 'F7.4' - E1: 'F7.4' = 'F7.4' mb')
26000      WRITE(8,700)A,B,AA,BB,W1,W2,WHT,WINDLIM,
26100      +      K,A0,A1,A2,A3,A4,A5,A6,RSCN,WINDCN,TEMCN
26200 700  FORMAT(/5X'PENMAN EQUATION COEFFICIENTS: '/5X'A='F6.3
26300      +' B='F6.3' A1='F6.3' B1='F6.3/5X'W1='F6.3' W2='F7.5
26400      +' WHT = 'F7.5' WINDLIM='F4.0//5X'RSO'I2': '4(E10.4,1X)/11X
26500      +3(E10.4,1X)//5X'RSCN='F7.5' WINDCN='F7.5' TEMCN='F7.5//)
26600      WMAY=SWMAY/31.
26700      WJUN=SWJUN/30.
26800      WJUL=SWJUL/31.
26900      WAUG=SWAUG/31.
27000      WSEP=SWSEP/30.
27100      AVWIND=(SWMAY+SWJUN+SWJUL+SWAUG+SWSEP)/153.
27200      WRITE(8,800)WMAY,WJUN,WJUL,WAUG,WSEP,AVWIND
27300 800  FORMAT(20X'W I N D'/5X'MAY: 'F6.1' MILES/DAY,
27400      +'5X'JUN: 'F6.1' MILES/DAY'/5X'JUL: 'F6.1' MILES/DAY,
27500      +'5X'AUG: 'F6.1' MILES/DAY'/5X'SEP: 'F6.1' MILES/DAY'/
27600      +'5X'AVERAGE SEASONAL WIND: 'F6.1' MILES/DAY')
27700      STOP
27800      END
27900
28000      SUBROUTINE SAVAPR(TF,SVP)

```

```

28100      TC=(TF-32.)*5./9.
28200      SVP= 1.3329*EXP(21.07 - 5336./(TC+273.1))
28300      RETURN
28400      END
28500
28600      SUBROUTINE JULDAY(IMN, IDAY, JD)
28700      DIMENSION MN(12)
28800      DATA MN/31,28,31,30,31,30,31,31,30,31,30,31/
28900      SUM=0
29000      DO 20 I=1,IMN
29100      IF (IMN.EQ.I)GOTO 10
29200      SUM=SUM+MN(I)
29300      10  JD=SUM+IDAY
29400      20  CONTINUE
29500      RETURN
29600      END
29700
29800      SUBROUTINE ROOZ(JD, IMN, IDAY)
29900      DIMENSION MN(12)
30000      DATA MN/31,28,31,30,31,30,31,31,30,31,30,31/
30100      SUM=0
30200      DO 10 J=1,12
30300      SUM=SUM+MN(J)
30400      IF ((JD-SUM).LE.0)GOTO 20
30500      10  CONTINUE
30600      20  IMN=J
30700      IDAY=JD - SUM + MN(J)
30800      RETURN
30900      END
31000      *
31100      *                END OF PROGRAM
31200      *
31300      *****

```

1800	LOGANTHYP	81781	53	43	48	49	44	35	83	090	
1801	LOGANTHYP	81791	53	43	48	49	44	35	83	048	
1802	LOGANTHYP	81801	63	53	58	59	54	29	319	071	
1803	LOGANTHYP	81811	74	63	68	70	65	22	872	164	
1804	LOGANTHYP	81821	84	74	79	81	76	16	282	045	
1805	LOGANTHYP	81831	94	84	89	91	86	9	180	253	
1806	LOGANTHYP	81841	73	63	68	70	65	7	637	107	
1807	LOGANTHYP	81851	78	68	73	75	70	12	711	347	
1808	LOGANTHYP	81861	78	68	73	75	70	5	743	303	
1809	LOGANTHYP	81871	88	78	83	85	80	14	400	162	
1810	LOGANTHYP	81881	79	69	74	76	71	101	708	407	
1811	LOGANTHYP	81891	78	68	73	75	70	14	780	118	
1812	LOGANTHYP	81901	78	68	73	75	70	20	406	234	
1813	LOGANTHYP	81911	69	59	64	66	61	7	529	139	
1814	LOGANTHYP	81921	74	64	69	71	66	16	732	266	
1815	LOGANTHYP	81931	80	70	75	77	72	10	769	269	
1816	LOGANTHYP	81941	81	71	76	78	73	21	634	299	
1817	LOGANTHYP	81951	81	71	76	78	73	19	372	179	
1818	LOGANTHYP	81961	84	74	79	81	76	39	620	332	
1819	LOGANTHYP	81971	65	55	60	62	57	04	30	625	106
1820	LOGANTHYP	81981	73	63	68	70	65	14	343	210	
1821	LOGANTHYP	81991	71	61	66	68	63	11	573	107	
1822	LOGANTHYP	82001	78	68	73	75	70	20	496	285	
1823	LOGANTHYP	82011	83	73	78	80	75	35	450	337	
1824	LOGANTHYP	82021	38	28	33	35	30	02	21	432	089
1825	LOGANTHYP	82031	46	36	41	43	38	12	818	110	
1826	LOGANTHYP	82041	82	72	77	79	74	41	782	313	
1827	LOGANTHYP	82051	70	60	65	67	62	42	726	339	
1828	LOGANTHYP	82061	76	66	71	73	68	27	309	216	
1829	LOGANTHYP	82071	80	70	75	77	72	31	670	281	
1830	LOGANTHYP	82081	83	73	78	80	75	34	798	381	
1831	LOGANTHYP	82091	82	72	77	79	74	37	766	422	
1832	LOGANTHYP	82101	87	77	82	84	79	35	802	217	
1833	LOGANTHYP	82111	89	79	84	86	81	17	790	293	
1834	LOGANTHYP	82121	86	76	81	83	78	33	784	222	
1835	LOGANTHYP	82131	82	72	77	79	74	10	914	217	
1836	LOGANTHYP	82141	87	77	82	84	79	28	122	337	





7100	LOGANUTNF	62781	90	72	64	83	31	139	247	
7200	LOGANUTNF	62881	85	70	57	79	10	794	204	
7300	LOGANUTNF	62981	85	54	62	82	9	782	276	
7400	LOGANUTNF	63081	90	50	64	88	7	723	232	
7500	LOGANUTNF	70181	85	62	62	74	23	438	218	
7600	LOGANUTNF	70281	81	56	62	79	03	34	546	270
7700	LOGANUTNF	70381	85	56	63	77	14	10	681	207
7800	LOGANUTNF	70481	89	72	64	88		41	469	329
7900	LOGANUTNF	70581	93	62	65	91		2	768	310
8000	LOGANUTNF	70681	83	62	64	84		16	626	319
8100	LOGANUTNF	70781	85	62	59	76	06	40	839	306
8200	LOGANUTNF	70881	78	41	55	78		20	791	215
8300	LOGANUTNF	70981	89	50	63	88		15	765	277
8400	LOGANUTNF	71081	89	63	67	87		38	687	377
8500	LOGANUTNF	71181	90	59	61	88	04	57	719	313
8600	LOGANUTNF	71281	90	54	62	88		19	727	216
8700	LOGANUTNF	71381	90	60	58	85		34	795	341
8800	LOGANUTNF	71481	86	47	58	84		15	771	241
8900	LOGANUTNF	71581	88	48	55	87		10	785	280
9000	LOGANUTNF	71681	91	49	60	90		17	649	269
9100	LOGANUTNF	71781	87	58	64	86		29	730	188
9200	LOGANUTNF	71881	88	52	61	87		17	573	363
9300	LOGANUTNF	71981	85	52	58	85		11	767	247
9400	LOGANUTNF	72081	89	49	61	88		28	761	322
9500	LOGANUTNF	72181	89	50	59	87		19	295	295
9600	LOGANUTNF	72281	89	49	58	86		21	751	345
9700	LOGANUTNF	72381	90	50	61	89		13	718	258
9800	LOGANUTNF	72481	90	50	61	85		16	659	224
9900	LOGANUTNF	72581	88	54	59	86		36	665	265
10000	LOGANUTNF	72681	72	60	55	68	25	34	424	434
10100	LOGANUTNF	72781	81	46	68	81		19	748	155
10200	LOGANUTNF	72881	88	48	59	86		15	735	278
10300	LOGANUTNF	72981	93	54	61	91		15	628	310
10400	LOGANUTNF	73081	88	58	60	87		8	695	333
10500	LOGANUTNF	73181	90	51	58	89		7	712	289
10600	LOGANUTNF	80181	92	49	63	90		19	717	402
10700	LOGANUTNF	80281	91	55	60	89		40	718	303
10800	LOGANUTNF	80381	90	50	58	88		17	716	280
10900	LOGANUTNF	80481	90	49	59	89		20	712	286
11000	LOGANUTNF	80581	92	54	60	90		14	714	319
11100	LOGANUTNF	80681	92	49	60	90		14	702	275
11200	LOGANUTNF	80781	91	59	62	90		44	685	315
11300	LOGANUTNF	80881	94	54	62	91		33	675	379
11400	LOGANUTNF	80981	86	64	62	85		126	314	456
11500	LOGANUTNF	81081	83	63	62	80	04	168	705	363
11600	LOGANUTNF	81181	86	49	60	85		65	605	264
11700	LOGANUTNF	81281	87	49	60	85		1	556	225
11800	LOGANUTNF	81381	74	50	56	73	03	3	319	280
11900	LOGANUTNF	81481	88	56	62	86		10	619	219
12000	LOGANUTNF	81581	88	54	58	70		31	578	366
12100	LOGANUTNF	81681	89	53	63	87		25	651	185
12200	LOGANUTNF	81781	91	51	61	89		4	660	259
12300	LOGANUTNF	81881	93	52	59	90		26	636	310
12400	LOGANUTNF	81981	86	58	60	86		44	622	378
12500	LOGANUTNF	82081	88	56	62	86		28	571	243
12600	LOGANUTNF	82181	88	55	59	80	26	39	645	326
12700	LOGANUTNF	82281	88	48	61	86		14	641	256
12800	LOGANUTNF	82381	91	47	61	88		11	634	232
12900	LOGANUTNF	82481	90	60	62	89		30	622	273
13000	LOGANUTNF	82581	93	58	62	91		16	615	256
13100	LOGANUTNF	82681	93	56	63	91		41	584	356
13200	LOGANUTNF	82781	92	55	57	78		48	518	333
13300	LOGANUTNF	82881	93	52	62	89		3	608	219
13400	LOGANUTNF	82981	91	58	62	88		44	571	326
13500	LOGANUTNF	83081	89	54	55	74		24	524	286
13600	LOGANUTNF	83181	78	42	51	76		11	621	188
13700	LOGANUTNF	90181	84	40	55	83		15	605	253
13800	LOGANUTNF	90281	80	53	54	79		14	518	308
13900	LOGANUTNF	90381	84	41	60	80		15	586	239
14000	LOGANUTNF	90481	88	47	60	87		15	563	101



14100	LOGANUTNF	90581	87	59	57	62	11	58	109	131
14200	LOGANUTNF	90681	87	62	56	53		21	213	200
14300	LOGANUTNF	90781	83	74	58	75		25	377	190
14400	LOGANUTNF	90881	81	78	61	78		24	508	158
14500	LOGANUTNF	90981	80	51	58	78		7	525	193
14600	LOGANUTNF	91081	82	58	61	81		16	540	202
14700	LOGANUTNF	91181	86	49	62	85		24	549	221
14800	LOGANUTNF	91281	87	50	59	77		6	541	255
14900	LOGANUTNF	91381	87	52	60	85		9	552	168
15000	LOGANUTNF	91481	86	49	58	80		11	545	214
15100	LOGANUTNF	91581	85	48	57	85		7	538	195
15200	LOGANUTNF	91681	87	49	58	86		16	531	243
15300	LOGANUTNF	91781	88	46	58	85		6	525	262
15400	LOGANUTNF	91881	88	47	59	86		15	520	201
15500	LOGANUTNF	91981	86	52	53	69		30	450	270
15600	LOGANUTNF	92081	78	41	54	76		14	516	201
15700	LOGANUTNF	92181	77	74	53	74		19	475	207
15800	LOGANUTNF	92281	77	73	52	73		14	423	198
15900	LOGANUTNF	92381	76	72	50	72		26	483	171
16000	LOGANUTNF	92481	79	50	55	75		24	482	182
16100	LOGANUTNF	92581	62	45	49	59	43	47	184	156
16200	LOGANUTNF	92681	63	32	47	62		13	349	078
16300	LOGANUTNF	92781	80	36	54	76		16	496	147
16400	LOGANUTNF	92881	78	51	54	75		60	448	227
16500	LOGANUTNF	92981	75	43	52	64		17	314	150
16600	LOGANUTNF	93081	64	38	46	62	05	22	486	134
16700	GREENVILLE LOGAN UT 1981									
16800	36.01=SEASONAL PENMAN ET									
16900	1.05									
17000	0.90									
17100	0.88									
17200	0.63									
17300	0.53									
17400	0.97									
17500	1.20									
17600	1.28									
17700	1.02									
17800	1.36									
17900	1.50									
18000	1.20									
18100	1.27									
18200	1.52									
18300	1.62									
18400	1.45									
18500	1.42									
18600	1.33									
18700	1.49									
18800	1.55									
18900	1.18									
19000	1.30									
19100	1.38									
19200	1.36									
19300	1.25									
19400	0.89									
19500	1.09									
19600	1.03									
19700	0.99									
19800	0.89									
19900	0.48									

```

100 *****
200 *
300 *   THIS PROGRAM ESTIMATES THE EXTRATERRESTRAL SOLAR RADIATION *
400 *   FOR THE GROWING SEASON  APRIL 1- SEPTEMBER 30  FOR THE LATI- *
500 *   TUDES OF THE STATE OF UTAH. *
600 *
700 *****
800 *
900   INTEGER RA(300,11)
1000  CHARACTER MON*3
1100  DIMENSION XLAT(11),K(300,11),MON(12)
1200  DATA XLAT/37.,37.5,38.,38.5,39.,39.5,40.,40.5,41.,41.5,42./
1300  DATA MON/'JAN','FEB','MAR','APR','MAY','JUN','JUL','AUG',
1400  +        'SEP','OCT','NOV','DEC'/
1500  DO 28 I=1,11
1600  PHI=XLAT(I)/57.2958
1700  DO 25 J=91,273
1800  D=J
1900  THETA=0.0172*(D-2.)
2000  IF(J.LT.3.)THETA=0.0172*(D+FLOAT(NDYR-2))
2100  RR=(1.+0.0167238*COS(THETA))/0.99986
2200  PC=0.0172*(D-1.)
2300  SIND=0.39785*SIN(PC+(279.9348+1.914827*SIN(PC)-0.079525*COS(PC)
2400  $+0.019938*SIN(2.*PC)-0.00162*COS(2.*PC))/57.29578)
2500  DER=ASIN(SIND)
2600  XSIN=SIN(PHI)*SIND
2700  XCOS=COS(PHI)*COS(DER)
2800  H=ACOS((-0.01454-XSIN)/XCOS)
2900  DL=7.63944*H
3000  RA(J,I)=118.5*RR*RR*(DL*XSIN+7.63944*XCOS*SIN(H))
3100  25 CONTINUE
3200  28 CONTINUE
3300  DOAR=49
3400  PAGE=0
3500  DO J=91,273
3600  CALL ROOZ(J,JMON,JDAY)
3700  IF((DOAR.EQ.49).AND.(PAGE.LT.2))THEN
3800  WRITE(90,35)
3900  DOAR=0
4000  PAGE=PAGE+1
4100  ENDIF
4200  IF((DOAR.EQ.50).AND.(PAGE.EQ.2))THEN
4300  WRITE(90,40)
4400  WRITE(90,35)
4500  DOAR=0
4600  ENDIF
4700  WRITE(90,30)MON(JMON),JDAY,(RA(J,I),I=1,11)
4800  DOAR=DOAR+1
4900  END DO
5000  30 FORMAT(14X,A3,I2,11(1X,I4))
5100  35 FORMAT(//14X,60(' '),//25X'EXTRATERRESTRIAL RADIATION (LANGLEYS
5200  +/DAY)'//42X'LATITUDE'/14X
5300  +'MON DAY 37 37.5 38 38.5 39 39.5 40 40.5 41 41.5 42'/
5400  +14X,60(' '))
5500  40 FORMAT(/)
5600  STOP
5700  END
5800  SUBROUTINE ROOZ(JA,JAMON,JADAY)
5900  DIMENSION JMON(12)
6000  DATA JMON/31,28,31,30,31,30,31,31,30,31,30,31/
6100  SMON=0
6200  DOAR=0
6300  DO K=1,12
6400  SMON=SMON+JMON(K)
6500  IF((SMON.GE.JA).AND.(DOAR.EQ.0))THEN
6600  JAMON=K
6700  JADAY=JA-(SMON-JMON(K))
6800  DOAR=1
6900  ENDIF
7000  ENDDO
7100  RETURN
7200  END

```

Appendix III: Isolines of Mean  
Longterm July Maximum and  
Minimum Temperatures for  
the State of Utah

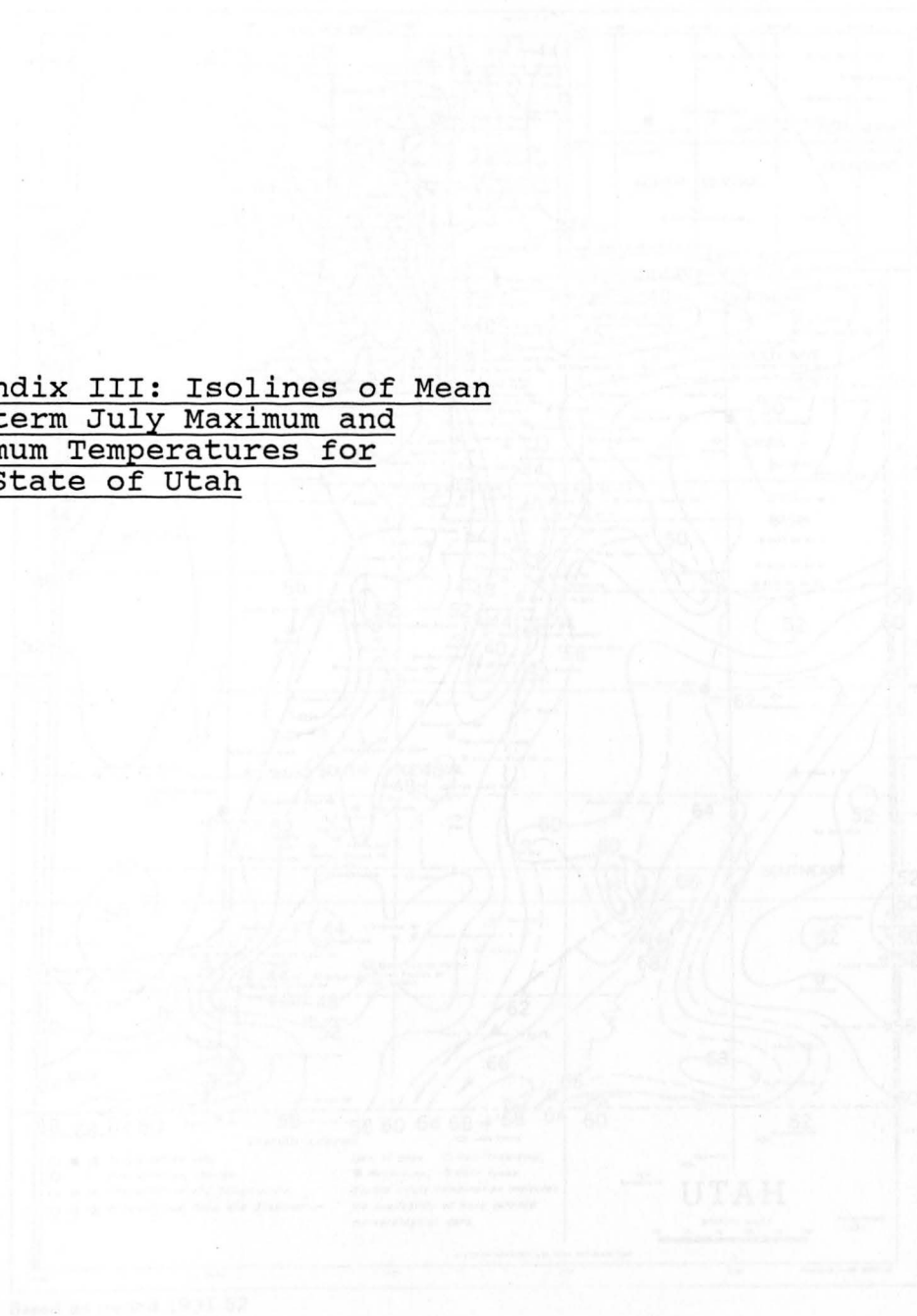
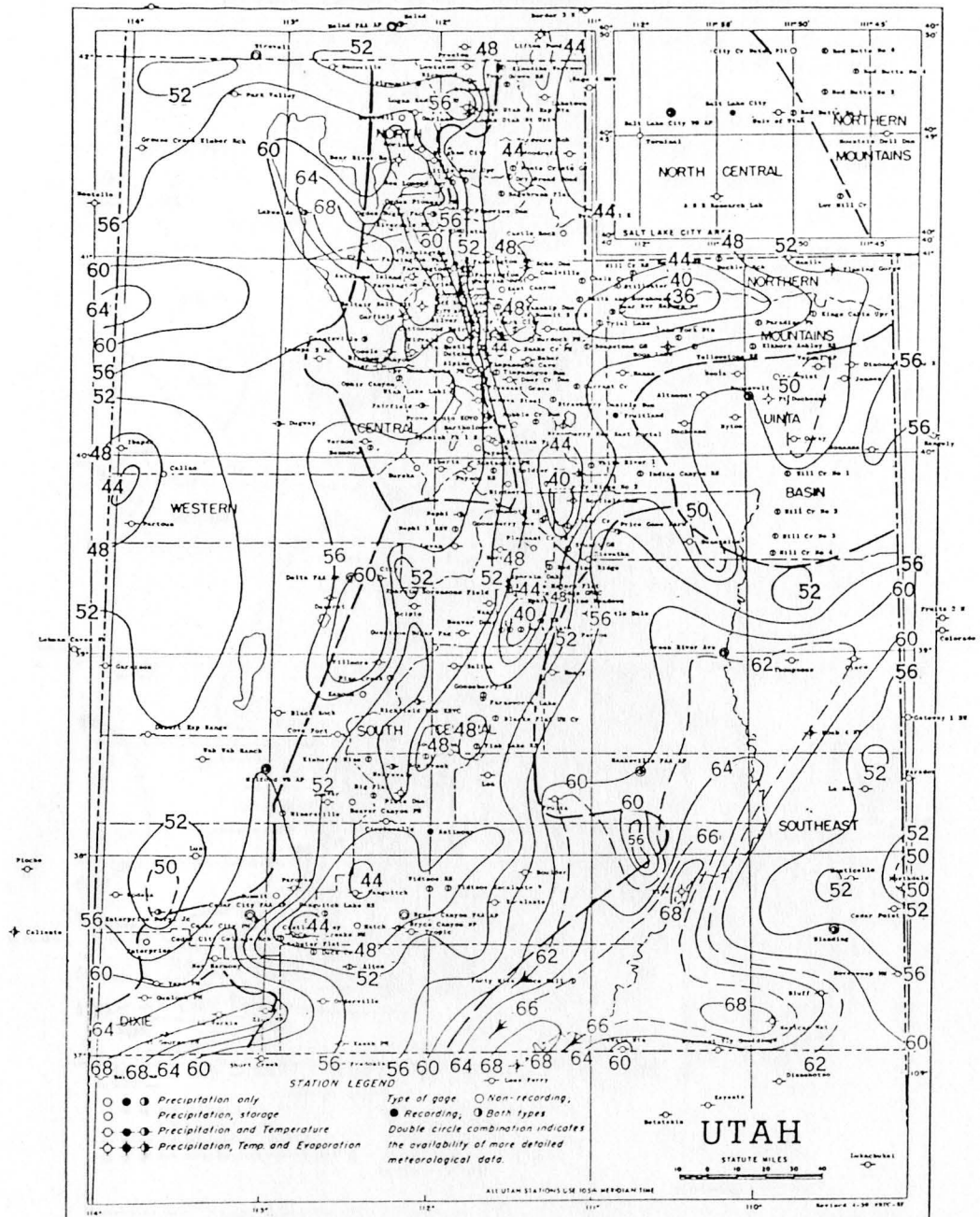


FIG. 13. - Mean Longterm July Minimum Temperature Isolines for the State of Utah. From (25).

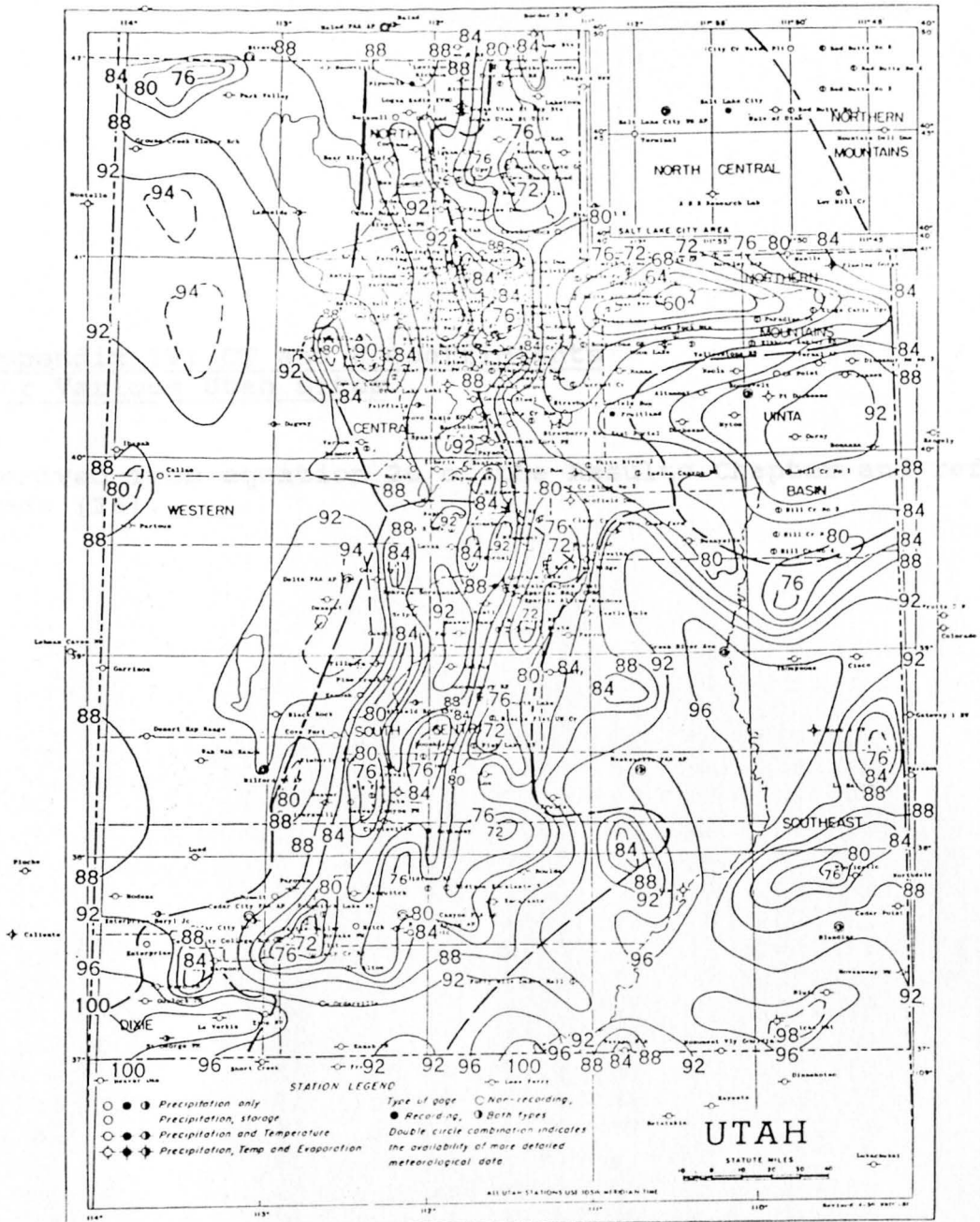
### Mean Minimum Temperature (°F.), July



Based on period 1931-52

FIG. 13.- Mean Longterm July Minimum Temperature Isolines for the State of Utah. From (25).

### Mean Maximum Temperature (°F.), July



Based on period 1931-52

FIG. 14.- Mean Longterm July Maximum Temperature Isolines for the State of Utah. From (25).



	LAT	LONG	ELEV	TXN	DM	K	CT
ALBANY	37.53	111.75	4930	50	83	.001267	.01137
ALBUQUERQUE	35.49	106.53	5250	52	82	.000917	.00773
ALBUQUERQUE	35.47	106.12	5270	52	84	.001073	.01101
ALBUQUERQUE	35.47	106.48	7040	51	84	.001077	.01113
ALBUQUERQUE	35.47	106.37	5230	50	83	.000913	.00783
ALBUQUERQUE	35.47	106.73	5710	50	82	.000904	.00772
ALBUQUERQUE	35.47	106.97	5890	51	83	.001015	.00775
ALBUQUERQUE	35.47	109.47	6110	54	85	.001174	.01117
ALBUQUERQUE	35.47	109.87	4710	54	83	.001180	.01120
ALBUQUERQUE	35.47	109.37	5450	54	83	.001074	.00782
ALBUQUERQUE	35.47	109.18	5910	53	82	.001071	.00782
ALBUQUERQUE	35.47	109.08	4760	54	88	.001187	.01122
ALBUQUERQUE	35.47	112.27	6060	43	84	.001200	.01124
ALBUQUERQUE	35.47	112.38	4620	50	84	.000944	.00910
ALBUQUERQUE	35.47	112.75	5250	52	83	.001090	.00783
ALBUQUERQUE	40.17	110.40	5510	53	81	.001173	.01103
ALBUQUERQUE	40.18	112.93	4340	54	83	.001075	.01014
ALBUQUERQUE	38.77	111.43	7640	52	82	.001012	.00911
ALBUQUERQUE	37.77	113.70	5300	57	90	.001133	.01092
ALBUQUERQUE	37.77	111.63	5810	53	84	.001311	.01218
ALBUQUERQUE	36.95	112.12	4480	56	82	.000854	.00786
ALBUQUERQUE	40.27	112.08	4880	56	82	.001106	.01013
ALBUQUERQUE	38.08	111.13	5930	57	83	.001167	.01164
ALBUQUERQUE	38.95	112.32	5120	60	92	.000894	.00814
ALBUQUERQUE	40.93	109.47	6270	52	84	.001098	.01021
ALBUQUERQUE	38.96	112.42	4700	60	90	.000943	.01089
ALBUQUERQUE	40.72	112.20	4330	62	85	.001143	.01148
ALBUQUERQUE	41.75	112.14	4400	58	87	.001081	.00981
ALBUQUERQUE	40.60	112.45	4230	61	90	.000987	.00958
ALBUQUERQUE	39.00	110.17	4970	62	96	.001099	.01142
ALBUQUERQUE	41.60	111.87	5860	45	76	.001384	.01250
ALBUQUERQUE	37.87	112.38	4000	58	96	.001183	.01231
ALBUQUERQUE	37.38	109.38	5240	54	93	.001137	.01293
ALBUQUERQUE	40.37	109.35	4700	57	92	.001100	.01117
ALBUQUERQUE	40.65	111.28	6470	47	80	.001122	.01043
ALBUQUERQUE	37.05	112.53	4950	56	93	.001158	.01120
ALBUQUERQUE	38.80	112.43	5010	55	88	.001200	.01118
ALBUQUERQUE	41.67	111.18	4260	61	92	.001075	.00986
ALBUQUERQUE	38.30	109.22	6720	56	88	.001103	.01036
ALBUQUERQUE	37.20	112.27	3230	68	100	.001199	.01099
ALBUQUERQUE	39.25	111.87	5315	56	88	.001098	.01013

Appendix IV: CT and K Coefficients  
for Various Utah Sites

Derived from equation 20 of the Results Chapter and reference (26).

SITE	LAT	LONG	ELEV	JULY		K	CT
				TMN	TMX		
ALPINE	40.45	111.78	4920	50	80	.001464	.01336
ALTA	40.60	111.63	8760	58	92	.000512	.00475
ALTAMONT	40.37	110.28	6370	52	84	.001079	.01001
ALTON	37.43	112.48	7040	51	84	.001077	.01012
BEAR RIV. REF.	41.47	112.27	4208	60	92	.000910	.00813
BIRDSEYE	39.87	111.53	5740	50	88	.001004	.00970
BLACK ROCK	38.70	112.95	4895	55	92	.001015	.00975
BLADING	37.62	109.47	6130	56	90	.001176	.01121
BLUFF	37.28	109.55	4315	64	97	.001343	.01272
BONANZA	40.02	109.18	5450	56	92	.001034	.00992
BRYCE CANYON	37.65	112.17	7915	48	80	.001071	.01002
CALLAO	39.90	113.72	4330	48	88	.001147	.01120
CASTLE DALE	39.22	111.02	5660	56	84	.001275	.01158
CEDAR POINT	37.72	109.08	6760	52	88	.001143	.01103
CIRCLEVILLE	38.17	112.27	6060	49	84	.001200	.01139
CORINNE	41.58	112.13	4240	57	92	.000897	.00831
DELTA	39.33	112.59	4623	60	94	.001069	.00923
DESERET	39.28	112.65	4585	56	94	.000944	.00915
DESERT EXP RN	38.60	113.75	5250	52	89	.001030	.00985
DUCHESNE	40.17	110.40	5510	52	86	.001175	.01105
DUGWAY	40.18	112.93	4340	54	90	.001076	.01016
EMERY	38.77	111.45	7640	52	82	.001012	.00932
ENTERPRIZE	37.57	113.70	5300	57	90	.001133	.01092
ESCALANTE	37.77	111.60	5810	53	84	.001331	.01238
EUREKA	39.95	112.12	6480	56	88	.000856	.00786
FAIRFIELD	40.27	112.08	4880	56	88	.001106	.01013
FERRON	39.08	111.13	5930	52	83	.001257	.01164
FILLMORE	38.95	112.32	5120	60	92	.000994	.00914
FLAMING GORGE	40.93	109.42	6270	52	84	.001099	.01021
FLOWELL	38.98	112.42	4702	60	90	.000998	.01069
GARFIELD	40.72	112.20	4330	62	85	.001349	.01148
GARLAND	41.75	112.19	4400	58	87	.001081	.00992
GRANTSVILLE	40.60	112.45	4290	61	90	.001087	.00958
GREEN RIVER	39.00	110.17	4070	62	96	.001209	.01142
HDWARE RNH	41.60	111.57	5560	45	76	.001364	.01250
HITE STN	37.87	110.38	4000	68	96	.001363	.01231
HOVENWEEP	37.38	109.08	5240	56	92	.001337	.01293
JENSEN	40.37	109.35	4760	52	92	.001130	.01117
KAMAS	40.65	111.28	6470	47	80	.001122	.01045
KANAB	37.05	112.53	4950	56	93	.001158	.01120
KANOSH	38.80	112.43	5010	55	88	.001200	.01118
KAYSVILLE	41.07	111.18	4267	61	92	.001075	.00986
LA SAL	38.30	109.22	6720	56	88	.001102	.01036
LA VERKIN	37.20	113.27	3220	68	100	.001199	.01099
LEVAN	39.55	111.87	5315	56	88	.001098	.01013

SITE	LAT	LONG	ELEV	JULY		K	CT
				TMN	TMX		
LOA	38.40	111.65	7080	51	80	.001172	.01077
LOGAN	41.75	111.82	4580	56	88	.001094	.00916
MANILA	41.00	109.73	6440	52	80	.001189	.01079
MANTI	39.25	111.63	5740	52	88	.001054	.01005
MARYSVALE	38.45	112.23	5910	52	86	.001134	.01068
MEXICAN HAT	37.15	109.87	4120	68	98	.001345	.01243
MOAB 4 NW	38.60	109.60	3965	62	96	.001343	.01275
MODENA	37.80	113.92	5480	52	90	.001008	.00975
MONTICELLO	37.87	109.30	6820	52	84	.001245	.01173
MONUMT VALY.	37.02	110.22	5300	63	94	.001202	.01122
MORONI	39.53	111.58	5525	51	92	.000906	.00902
NEOLA	40.42	110.03	5920	51	82	.001248	.01153
NEPHI	39.70	111.83	5130	56	92	.000947	.00901
NEW HRMNY	37.48	113.30	5290	60	88	.001210	.01091
OAK CITY	39.38	112.33	5070	60	94	.000869	.00810
OURAY 4 NE	40.13	109.65	4670	50	92	.001152	.01157
PANGUITCH	37.82	112.45	6720	44	84	.001096	.01072
PARADISE	41.58	111.60	5000	53	83	.001073	.01026
PARK CITY	40.72	111.52	6740	41	79	.001038	.01045
PARK VALY	41.82	113.33	5530	56	85	.000877	.00766
PAROWAN	37.85	112.83	5930	56	88	.001161	.00966
PARTOUN	39.63	113.88	4780	47	88	.001054	.01038
PAYSON	40.03	111.72	4800	56	89	.001128	.01045
PINE VW DAM	41.25	111.83	4940	52	88	.001005	.00948
PLEASANT GV	40.37	111.72	4760	56	88	.001150	.01055
RANDOLPH	41.75	111.13	6280	44	81	.001074	.01017
RICHMOND	41.90	111.82	4680	52	88	.000991	.00932
ROOSEVELT	40.30	109.98	5104	52	89	.001150	.01107
SLC AIRPORT	40.78	111.95	4267	61	92	.001032	.00869
ST. GEORGE	37.08	113.68	2800	69	101	.001057	.01054
SALINA	38.97	111.87	5130	52	92	.001014	.01001
SNOWVILLE	41.97	112.72	4560	52	88	.000940	.00878
SUMMIT	37.80	112.93	6000	56	88	.001058	.00982
SUNNYSIDE	39.57	110.37	6780	51	88	.000934	.00899
THOMPSON	38.97	109.72	5150	62	96	.000990	.00937
THORNACK	41.75	111.13	6280	44	81	.001080	.01035
TIMPANOGAS	40.45	111.70	5640	52	88	.000961	.00911
TOOELE	40.53	112.30	5070	60	90	.000936	.00835
TROPIC	37.63	112.08	6280	50	83	.001249	.01174
UTAH LAKE	40.37	111.90	4497	56	84	.001406	.01257
VERNON	40.08	112.45	5485	56	88	.000977	.00894
WAH WAH RN	38.48	113.42	4880	54	90	.001099	.01045
WOODRUFF	41.53	111.15	6315	44	80	.001084	.01025

EXTRATERRESTRIAL RADIATION (LANGLEYS/HR)

	LATITUDE										
	35	35.5	36	36.5	37	37.5	38	38.5	39	39.5	40
APR 1	740	741	742	743	744	745	746	747	748	749	750
APR 2	740	741	742	743	744	745	746	747	748	749	750
APR 3	740	741	742	743	744	745	746	747	748	749	750
APR 4	740	741	742	743	744	745	746	747	748	749	750
APR 5	740	741	742	743	744	745	746	747	748	749	750
APR 6	740	741	742	743	744	745	746	747	748	749	750
APR 7	740	741	742	743	744	745	746	747	748	749	750
APR 8	740	741	742	743	744	745	746	747	748	749	750
APR 9	740	741	742	743	744	745	746	747	748	749	750
APR 10	740	741	742	743	744	745	746	747	748	749	750
APR 11	740	741	742	743	744	745	746	747	748	749	750
APR 12	740	741	742	743	744	745	746	747	748	749	750
APR 13	740	741	742	743	744	745	746	747	748	749	750
APR 14	740	741	742	743	744	745	746	747	748	749	750
APR 15	740	741	742	743	744	745	746	747	748	749	750
APR 16	740	741	742	743	744	745	746	747	748	749	750
APR 17	740	741	742	743	744	745	746	747	748	749	750
APR 18	740	741	742	743	744	745	746	747	748	749	750
APR 19	740	741	742	743	744	745	746	747	748	749	750
APR 20	740	741	742	743	744	745	746	747	748	749	750
APR 21	740	741	742	743	744	745	746	747	748	749	750
APR 22	740	741	742	743	744	745	746	747	748	749	750
APR 23	740	741	742	743	744	745	746	747	748	749	750
APR 24	740	741	742	743	744	745	746	747	748	749	750
APR 25	740	741	742	743	744	745	746	747	748	749	750
APR 26	740	741	742	743	744	745	746	747	748	749	750
APR 27	740	741	742	743	744	745	746	747	748	749	750
APR 28	740	741	742	743	744	745	746	747	748	749	750
APR 29	740	741	742	743	744	745	746	747	748	749	750
APR 30	740	741	742	743	744	745	746	747	748	749	750
MAY 1	740	741	742	743	744	745	746	747	748	749	750
MAY 2	740	741	742	743	744	745	746	747	748	749	750
MAY 3	740	741	742	743	744	745	746	747	748	749	750
MAY 4	740	741	742	743	744	745	746	747	748	749	750
MAY 5	740	741	742	743	744	745	746	747	748	749	750
MAY 6	740	741	742	743	744	745	746	747	748	749	750
MAY 7	740	741	742	743	744	745	746	747	748	749	750
MAY 8	740	741	742	743	744	745	746	747	748	749	750
MAY 9	740	741	742	743	744	745	746	747	748	749	750
MAY 10	740	741	742	743	744	745	746	747	748	749	750
MAY 11	740	741	742	743	744	745	746	747	748	749	750
MAY 12	740	741	742	743	744	745	746	747	748	749	750
MAY 13	740	741	742	743	744	745	746	747	748	749	750
MAY 14	740	741	742	743	744	745	746	747	748	749	750
MAY 15	740	741	742	743	744	745	746	747	748	749	750
MAY 16	740	741	742	743	744	745	746	747	748	749	750

Appendix V: Estimated Extra-terrestrial Solar Radiation for Utah Latitudes

Derived from program EXTRAD listed in Appendix II.

## EXTRATERRESTRIAL RADIATION (LANGLEYS/DAY)

MON DAY	LATITUDE										
	37	37.5	38	38.5	39	39.5	40	40.5	41	41.5	42
APR 1	785	781	777	773	769	764	760	756	751	747	742
APR 2	790	786	782	778	774	770	766	761	757	753	748
APR 3	795	791	788	784	780	775	771	767	763	758	754
APR 4	800	797	793	789	785	781	777	773	768	764	760
APR 5	805	802	798	794	790	786	782	778	774	770	766
APR 6	810	807	803	799	796	792	788	784	780	776	771
APR 7	815	812	808	804	801	797	793	789	785	781	777
APR 8	820	817	813	810	806	802	798	795	791	787	783
APR 9	825	822	818	815	811	807	804	800	796	792	788
APR10	830	826	823	820	816	813	809	805	802	798	794
APR11	835	831	828	825	821	818	814	811	807	803	799
APR12	839	836	833	830	826	823	819	816	812	809	805
APR13	844	841	838	834	831	828	824	821	818	814	810
APR14	848	845	842	839	836	833	829	826	823	819	816
APR15	853	850	847	844	841	838	834	831	828	824	821
APR16	857	854	852	849	846	843	839	836	833	830	826
APR17	862	859	856	853	850	847	844	841	838	835	831
APR18	866	863	861	858	855	852	849	846	843	840	837
APR19	870	868	865	862	860	857	854	851	848	845	842
APR20	874	872	869	867	864	861	858	856	853	850	847
APR21	879	876	874	871	869	866	863	860	857	855	852
APR22	883	880	878	875	873	870	868	865	862	859	856
APR23	887	884	882	880	877	875	872	870	867	864	861
APR24	891	889	886	884	882	879	877	874	871	869	866
APR25	895	893	890	888	886	883	881	879	876	873	871
APR26	898	896	894	892	890	888	885	883	880	878	875
APR27	902	900	898	896	894	892	890	887	885	882	880
APR28	906	904	902	900	898	896	894	891	889	887	884
APR29	910	908	906	904	902	900	898	896	893	891	889
APR30	913	911	910	908	906	904	902	900	898	895	893
MAY 1	917	915	913	912	910	908	906	904	902	900	897
MAY 2	920	919	917	915	913	912	910	908	906	904	902
MAY 3	924	922	920	919	917	915	914	912	910	908	906
MAY 4	927	925	924	922	921	919	917	916	914	912	910
MAY 5	930	929	927	926	924	923	921	919	918	916	914
MAY 6	933	932	931	929	928	926	925	923	921	920	918
MAY 7	936	935	934	933	931	930	928	927	925	923	922
MAY 8	940	938	937	936	935	933	932	930	929	927	925
MAY 9	943	941	940	939	938	937	935	934	932	931	929
MAY10	945	944	943	942	941	940	938	937	936	934	933
MAY11	948	947	946	945	944	943	942	940	939	938	936
MAY12	951	950	949	948	947	946	945	944	942	941	940
MAY13	954	953	952	951	950	949	948	947	946	944	943
MAY14	956	956	955	954	953	952	951	950	949	948	946
MAY15	959	958	958	957	956	955	954	953	952	951	950
MAY16	962	961	960	959	959	958	957	956	955	954	953



## EXTRATERRESTRIAL RADIATION (LANGLEYS/DAY)

MON DAY	LATITUDE										
	37	37.5	38	38.5	39	39.5	40	40.5	41	41.5	42
MAY17	964	963	963	962	961	961	960	959	958	957	956
MAY18	966	966	965	965	964	963	962	962	961	960	959
MAY19	969	968	968	967	967	966	965	964	964	963	962
MAY20	971	970	970	970	969	968	968	967	966	965	964
MAY21	973	973	972	972	971	971	970	970	969	968	967
MAY22	975	975	975	974	974	973	973	972	971	971	970
MAY23	977	977	977	976	976	975	975	974	974	973	972
MAY24	979	979	979	978	978	978	977	977	976	976	975
MAY25	981	981	981	980	980	980	979	979	978	978	977
MAY26	983	983	983	982	982	982	982	981	981	980	980
MAY27	985	985	984	984	984	984	984	983	983	982	982
MAY28	986	986	986	986	986	986	986	985	985	984	984
MAY29	988	988	988	988	988	988	987	987	987	986	986
MAY30	989	989	990	990	989	989	989	989	989	988	988
MAY31	991	991	991	991	991	991	991	991	991	990	990
JUN 1	992	992	993	993	993	993	993	992	992	992	992
JUN 2	994	994	994	994	994	994	994	994	994	994	993
JUN 3	995	995	995	995	996	996	996	996	995	995	995
JUN 4	996	996	997	997	997	997	997	997	997	997	997
JUN 5	997	997	998	998	998	998	998	998	998	998	998
JUN 6	998	999	999	999	999	999	1000	1000	1000	1000	999
JUN 7	999	1000	1000	1000	1000	1001	1001	1001	1001	1001	1001
JUN 8	1000	1000	1001	1001	1001	1002	1002	1002	1002	1002	1002
JUN 9	1001	1001	1002	1002	1002	1003	1003	1003	1003	1003	1003
JUN10	1002	1002	1003	1003	1003	1003	1004	1004	1004	1004	1004
JUN11	1002	1003	1003	1004	1004	1004	1004	1005	1005	1005	1005
JUN12	1003	1003	1004	1004	1005	1005	1005	1005	1006	1006	1006
JUN13	1004	1004	1005	1005	1005	1006	1006	1006	1006	1006	1006
JUN14	1004	1005	1005	1005	1006	1006	1006	1007	1007	1007	1007
JUN15	1004	1005	1005	1006	1006	1007	1007	1007	1007	1008	1008
JUN16	1005	1005	1006	1006	1007	1007	1007	1008	1008	1008	1008
JUN17	1005	1006	1006	1007	1007	1007	1008	1008	1008	1008	1009
JUN18	1005	1006	1006	1007	1007	1008	1008	1008	1009	1009	1009
JUN19	1005	1006	1007	1007	1007	1008	1008	1008	1009	1009	1009
JUN20	1005	1006	1007	1007	1008	1008	1008	1009	1009	1009	1009
JUN21	1005	1006	1007	1007	1008	1008	1008	1009	1009	1009	1009
JUN22	1005	1006	1007	1007	1007	1008	1008	1009	1009	1009	1009
JUN23	1005	1006	1006	1007	1007	1008	1008	1008	1009	1009	1009
JUN24	1005	1006	1006	1007	1007	1008	1008	1008	1008	1009	1009
JUN25	1005	1005	1006	1006	1007	1007	1008	1008	1008	1008	1008
JUN26	1004	1005	1006	1006	1006	1007	1007	1007	1008	1008	1008
JUN27	1004	1005	1005	1006	1006	1006	1007	1007	1007	1007	1008
JUN28	1003	1004	1005	1005	1005	1006	1006	1006	1007	1007	1007
JUN29	1003	1004	1004	1004	1005	1005	1006	1006	1006	1006	1006
JUN30	1002	1003	1003	1004	1004	1005	1005	1005	1005	1005	1006
JUL 1	1002	1002	1003	1003	1003	1004	1004	1004	1004	1005	1005
JUL 2	1001	1001	1002	1002	1003	1003	1003	1003	1004	1004	1004

## EXTRATERRESTRIAL RADIATION (LANGLEYS/DAY)

MON DAY	LATITUDE										
	37	37.5	38	38.5	39	39.5	40	40.5	41	41.5	42
JUL 3	1000	1001	1001	1001	1002	1002	1002	1002	1003	1003	1003
JUL 4	999	1000	1000	1000	1001	1001	1001	1001	1001	1002	1002
JUL 5	998	999	999	999	1000	1000	1000	1000	1000	1000	1000
JUL 6	997	998	998	998	999	999	999	999	999	999	999
JUL 7	996	996	997	997	997	998	998	998	998	998	998
JUL 8	995	995	996	996	996	996	996	996	996	996	996
JUL 9	994	994	994	995	995	995	995	995	995	995	995
JUL10	992	993	993	993	993	993	993	993	993	993	993
JUL11	991	991	991	992	992	992	992	992	992	992	991
JUL12	990	990	990	990	990	990	990	990	990	990	990
JUL13	988	988	988	989	989	989	988	988	988	988	988
JUL14	987	987	987	987	987	987	987	987	986	986	986
JUL15	985	985	985	985	985	985	985	985	984	984	984
JUL16	983	983	983	983	983	983	983	983	982	982	982
JUL17	981	981	981	981	981	981	981	981	980	980	979
JUL18	980	980	980	979	979	979	979	978	978	978	977
JUL19	978	978	978	977	977	977	977	976	976	975	975
JUL20	976	976	975	975	975	975	974	974	973	973	972
JUL21	974	974	973	973	973	972	972	971	971	970	970
JUL22	972	971	971	971	970	970	970	969	968	968	967
JUL23	969	969	969	969	968	968	967	967	966	965	964
JUL24	967	967	967	966	966	965	965	964	963	962	962
JUL25	965	965	964	964	963	963	962	961	960	960	959
JUL26	963	962	962	961	961	960	959	958	958	957	956
JUL27	960	960	959	959	958	957	956	956	955	954	953
JUL28	958	957	957	956	955	954	954	953	952	951	950
JUL29	955	955	954	953	952	952	951	950	949	948	947
JUL30	953	952	951	950	950	949	948	947	946	945	943
JUL31	950	949	948	948	947	946	945	944	943	941	940
AUG 1	947	946	946	945	944	943	942	940	939	938	937
AUG 2	944	944	943	942	941	940	938	937	936	935	933
AUG 3	941	941	940	939	938	936	935	934	933	931	930
AUG 4	939	938	937	935	934	933	932	931	929	928	926
AUG 5	936	935	933	932	931	930	928	927	926	924	923
AUG 6	933	931	930	929	928	926	925	924	922	921	919
AUG 7	929	928	927	926	924	923	922	920	918	917	915
AUG 8	926	925	924	922	921	919	918	916	915	913	911
AUG 9	923	922	920	919	917	916	914	913	911	909	907
AUG10	920	918	917	915	914	912	911	909	907	905	903
AUG11	916	915	913	912	910	909	907	905	903	901	899
AUG12	913	911	910	908	907	905	903	901	899	897	895
AUG13	910	908	906	905	903	901	899	897	895	893	891
AUG14	906	904	903	901	899	897	895	893	891	889	887
AUG15	902	901	899	897	895	893	891	889	887	885	882
AUG16	899	897	895	893	891	889	887	885	883	880	878
AUG17	895	893	891	889	887	885	883	881	878	876	873
AUG18	891	889	887	885	883	881	879	876	874	871	869

## EXTRATERRESTRIAL RADIATION (LANGLEYS/DAY)

MON DAY	LATITUDE										
	37	37.5	38	38.5	39	39.5	40	40.5	41	41.5	42
AUG19	887	885	883	881	879	877	874	872	869	867	864
AUG20	884	881	879	877	875	872	870	867	865	862	860
AUG21	880	877	875	873	871	868	866	863	860	858	855
AUG22	876	873	871	869	866	864	861	858	856	853	850
AUG23	872	869	867	864	862	859	857	854	851	848	845
AUG24	868	865	863	860	857	855	852	849	846	844	841
AUG25	863	861	858	856	853	850	847	845	842	839	836
AUG26	859	857	854	851	848	846	843	840	837	834	831
AUG27	855	852	850	847	844	841	838	835	832	829	826
AUG28	851	848	845	842	839	836	833	830	827	824	821
AUG29	846	843	841	838	835	832	828	825	822	819	815
AUG30	842	839	836	833	830	827	824	820	817	814	810
AUG31	837	834	831	828	825	822	819	815	812	808	805
SEP 1	833	830	827	824	820	817	814	810	807	803	800
SEP 2	828	825	822	819	815	812	809	805	802	798	794
SEP 3	824	821	817	814	811	807	804	800	796	793	789
SEP 4	819	816	812	809	806	802	798	795	791	787	784
SEP 5	814	811	808	804	801	797	793	790	786	782	778
SEP 6	810	806	803	799	796	792	788	784	780	777	773
SEP 7	805	801	798	794	790	787	783	779	775	771	767
SEP 8	800	796	793	789	785	781	778	774	770	766	761
SEP 9	795	791	788	784	780	776	772	768	764	760	756
SEP10	790	786	783	779	775	771	767	763	759	754	750
SEP11	785	781	778	774	770	766	761	757	753	749	744
SEP12	780	776	772	768	764	760	756	752	747	743	739
SEP13	775	771	767	763	759	755	751	746	742	737	733
SEP14	770	766	762	758	754	749	745	741	736	732	727
SEP15	765	761	757	753	748	744	739	735	730	726	721
SEP16	760	756	751	747	743	738	734	729	725	720	715
SEP17	755	750	746	742	737	733	728	724	719	714	709
SEP18	750	745	741	736	732	727	723	718	713	708	704
SEP19	744	740	735	731	726	722	717	712	707	703	698
SEP20	739	735	730	725	721	716	711	706	702	697	692
SEP21	734	729	725	720	715	710	706	701	696	691	686
SEP22	728	724	719	714	710	705	700	695	690	685	680
SEP23	723	718	714	709	704	699	694	689	684	679	674
SEP24	718	713	708	703	698	693	688	683	678	673	668
SEP25	712	707	703	698	693	688	682	677	672	667	662
SEP26	707	702	697	692	687	682	677	671	666	661	655
SEP27	701	696	691	686	681	676	671	666	660	655	649
SEP28	696	691	686	681	676	670	665	660	654	649	643
SEP29	691	685	680	675	670	664	659	654	648	643	637
SEP30	685	680	675	669	664	659	653	648	642	637	631