



1 **A MID-LATITUDE MAASTRICHTIAN PALYNOLOGICAL RECORD FROM**
2 **THE YACORAITE FORMATION (SALTA GROUP), NORTHWESTERN**
3 **ARGENTINA**

4 **Magdalena Llorens^{a, b, *}, Valeria S. Pérez Loinaze^{a, c}, Paula L. Narváez^{a, d}, Alfredo**
5 **M. Zelaya^e, Egly Pérez Pincheira^f, Sergio Gorustovich^e**

6 ^a Consejo Nacional de Investigaciones Científicas y Técnicas (CONICET), Argentina

7 ^b Laboratorio de Botánica and Laboratorio de Palinología, Facultad de Ciencias
8 Naturales, Universidad Nacional de la Patagonia San Juan Bosco, 9 de Julio 25,
9 (U9100CKN) Trelew, Chubut, Argentina

10 ^c Museo Argentino de Ciencias Naturales “B. Rivadavia”, Av. Ángel Gallardo 470
11 (C1405DJR), Buenos Aires, Argentina

12 ^d IANIGLA/CCT Mendoza. Av. Adrián Ruiz Leal s/n, Parque Gral. San Martín, (5500)
13 Mendoza, Argentina

14 ^e Comisión Nacional de Energía Atómica, Gerencia Exploración de Materias Primas,
15 Avenida del Libertador 8250, (1429) Ciudad Autónoma de Buenos Aires, Argentina

16 ^f Cátedra de Palinología, Facultad de Ciencias Naturales y Museo, Universidad
17 Nacional de La Plata, Calle 64, 3 (B1904DZA) La Plata, Buenos Aires, Argentina.

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19 *Corresponding author: Magdalena Llorens

20 E-mail address: magdalena.llorens@gmail.com

21

22 **Abstract.**

23 A palynological assemblage from one sample of the Yacoraite Formation outcrops
24 (Maastrichtian–Danian of the Salta Basin) at the Valle del Tonco, Salta Province,
25 Argentina, is described. The flora is distinctly continental and dominated by

26 angiosperms, accompanying ferns and gymnosperms, along with few freshwater algae,
27 bryophytes and fungi. The presence of taxa such as *Aquillapollenites magnus*,
28 *Gabonisporis vigorouxii* and *Crassitricolporites brasiliensis*, together with the absence
29 of typical characteristic taxa from of the Danian *Mtchedlishvilia saltenia* Zone, allow to
30 propose a late Maastrichtian age for this palynoflora. Due to its location at mid-
31 latitudes, the palynoflora from the Yacoraite Formation is within a transitional or mixed
32 flora zone with closer affinity to the equatorial Palmae Province. Therefore, the
33 recorded flora has ~~represents~~ a good potential opportunity to understand the transition
34 between the floras of low and high latitudes and to know the floristic composition of the
35 area before the K-Pg extinction event.

36 **Keywords.** Palynology; Late Cretaceous; Yacoraite Formation; Salta Group; Argentina.

37

38 1. Introduction

39 The Maastrichtian floras of South America are very well known at high latitudes (e.g.,
40 Clyde et al., 2021; Pérez Pincheira and Di Pasquo, 2021 and citations therein). On the
41 other hand, coeval low-latitude floras are known in broad temporal contexts, with data
42 obtained from borehole samples, mainly of marine environments and with a
43 biostratigraphic approach (e.g., Regali et al., 1984 a, b; Muller et al 1987). However,
44 there are very few records of palynofloras from mid-latitude regions (Pérez Leytón,
45 1987; Limarino et al., 2000; Quattrocchio et al., 2005).

46 In Argentina, the main unit with mid-latitude Cretaceous sediments is the Salta Group.

47 It is a geological unit at the Northwest of Argentina widely extended that includes strata
48 from the Lower Cretaceous to the Eocene (Salfity and Marquillas, 1994; Marquillas et
49 al., 2005). It has economic interest due to its hydrocarbon and uranium content, but also
50 its paleontological content is very relevant.

51 The fossil content of the Yacoraite Formation includes marine and continental biotas,
52 and has been analyzed and discussed in several works. Fossil fishes (Benedetto and
53 Sánchez, 1972; Cione, 1977; Cione et al., 1985; Cione and Pereyra, 1985), reptiles
54 (Gasparini and Buffetaut, 1980), gastropods (Cónsole-Gonella, 2011 and citations
55 therein), stromatolites (Villafañe et al., 2021), ostracods, scarce miliolid foraminifera
56 (Kielbowicz de Stach and Angelozzi, 1984; and Méndez and Viviers, 1973), carophytes
57 (Musacchio, 1972) and other algae, in addition to footprints of dinosaurs and birds
58 (Alonso, 1980; Alonso and Marquillas, 1986; Marquillas et al., 2005; Marquillas et al.,
59 2007; Marquillas and Salfity, 1994; de Valais and Cónsole-Gonella, 2018; Cónsole-
60 Gonella et al., 2021), have been recorded.

61 Nevertheless, the scarcity of floristic records in the Yacoraite Formation is notorious.
62 There are no reports of leaves or wood remains and the calcareous composition of the
63 rocks in this formation makes it difficult for palynomorph preservation. Therefore, the
64 palynomorph record is restricted to three preliminary reports. The first was made by
65 Moroni (1982), who compared the palynological content from the Yacoraite and
66 Olmedo formations at the Lomas de Olmedo sub-basin. Seven angiosperm taxa and the
67 presence of polylicate pollen grains were mentioned along with algae spores, but none
68 were described or illustrated. Another work is a conference abstract where the finding of
69 massulae of *Azolla cretacea* and dinoflagellates cysts is mentioned (Papú and Melendi,
70 1984). Finally, Narváez (2009) found a single sample from the Yacoraite Formation
71 with palynological evidence limited to three types of freshwater algae (*Leiosphaeridia*
72 sp. and *Pediastrum* spp.).

73 The aim of this contribution is to describe and illustrate for first time a palynological
74 assemblage from the Yacoraite Formation at the Don Otto uranium mine. A comparison

75 with other Late Cretaceous and Danian mid-latitude palynofloras, biostratigraphy and
76 discussion of the age and paleoecological inferences are also included.

77

78 **2. Geological setting**

79 The Salta Basin in northwestern Argentina is relevant for its wide extent, and
80 hydrocarbon production (Turic et al., 1987; Marquillas and Salfity, 1988; Gómez Omil
81 et al., 1989; Salfity and Marquillas, 1994). The basin fill corresponds to the Salta Group
82 (Turner, 19589) deposited from the Early Cretaceous up to the Eocene and composed of
83 three main units from base to top: Pirgua (Reyes and Salfity, 1973), Balbuena and Santa
84 Bárbara subgroups (Moreno, 1970). It reaches 5000 m thick and crops out in seven sub-
85 basins: Tres Cruces, Lomas de Olmedo, Alemanía, Metán, El Rey, Sey and Brealito
86 (Marquillas et al., 2005 and citations therein).

87 During the Maastrichtian, two major factors controlled the Salta Basin fill, the decrease
88 in tectonic subsidence and a relative sea level rise that affected South America
89 (Marquillas et al., 2005). In northern Argentina, an extensive and shallow Atlantic
90 marine ingression installed a carbonate system coincident with mainly humid conditions
91 until the Danian (Marquillas et al., 2005). Due to its calcareous-dolomitic composition
92 with an intense yellow of weathering color, the Yacoraite Formation (Groeber, 1952)
93 constitutes an excellent marker horizon in the Salta Group (Marquillas et al., 2005). The
94 Cretaceous–Paleogene transition is also contained in the Yacoraite Formation of
95 Maastrichtian–Danian age (Sial et al., 2001; Marquillas et al., 2003).

96 The main lithology of the Yacoraite Formation are carbonate-calcareous rocks and
97 dolomitic limestones, it also contains shales and sandstones, reaching a maximum
98 thickness of 200 m (Marquillas et al., 2007). In the Alemanía sub-basin, the Yacoraite
99 Formation (Balbuena Subgroup) crops out at the Valle del Tonco locality (Tonco River

100 valley), where it overlies the Lecho Formation (Balbuena Subgroup) and underlies the
101 Mealla Formation (Santa Bárbara Subgroup) within the Salta Group (Payrola et al.,
102 2011). At the Don Otto mine vicinity (Fig. 1.A–E), the unit crops out with a thickness
103 of 174 m mainly corresponding to the Don Otto Complex Member where Zelaya et al.
104 (2014a, b) recognized eight sedimentary facies (F1–8) grouped into four facies
105 associations (FA1–4) related to intertidal–subtidal deposits and lagoon-type
106 environments with sporadic storm events and subaerial exposure. The palynological
107 sample here studied was recovered from dark gray to black siltstones and claystones
108 (facies F8) within FA2.

109 **2.1. Age**

110 The widely recorded dinosaur footprints in the Yacoraite Formation (e.g., Cónsole-
111 Gonella et al., 2021) undoubtedly indicates a Cretaceous age for the unit. As
112 aforementioned, palynological studies from well samples in the Lomas de Olmedo sub-
113 basin, showed that the Yacoraite Formation finished its sedimentation in the Danian
114 (Moroni, 1982, 1984). This is consistent with the interpretation of the Yacoraite
115 Formation in the context of the transgressions that occurred in southern South America
116 during the Maastrichtian–Paleocene (Marquillas, 1985; Marquillas and Salfity, 1988;
117 Marquillas et al., 2005).

118 Despite having been the subject of diverse reviews (e.g., Castaños et al., 1975;
119 Marquillas, 1985; Gómez Omil et al., 1989; Salfity and Marquillas, 1994; Marquillas et
120 al., 2005), isotopic studies of this rock unit are still scarce (Matthews et al., 1996; Sial et
121 al., 2001; Marquillas et al., 2007). An absolute U-Pb zircon dating has confirmed a
122 Maastrichtian age for at least part of the unit, with a result of 71.9 ± 0.4 Ma and $68.4 \pm$
123 0.7 Ma from two tuff layers obtained from the Amblayo and Alemania members of the
124 Yacoraite Formation (Marquillas et al., 2011).

125

126 **3. Materials and methods**

127 Two samples were recovered from outcrops of the Yacoraite Formation in the Don Otto
128 mine, located at the southernmost section of the Valle del Tonco (Tonco River valley;
129 25°36'39"S, 65°55'10.6"W) (Fig. 1A–E). Two levels of grey-black pelitic sedimentary
130 rocks with high content of organic matter were sampled in the vicinity of the camp, but
131 only one contained palynomorphs (CNS-M 001/A–H; Fig. 2). The samples were treated
132 separately with hydrochloric and hydrofluoric acids following the palynological
133 techniques for extraction and concentration of palynomorphs by Phipps and Playford
134 (1984). The organic residues were sieved using 20µm and 10µm opening mesh and
135 mounted in permanent slides. Observations were made with an Olympus BX-51
136 microscope. Photographs were taken with a Nikon DS-Fi1 digital camera. Reference to
137 the illustrated specimens is given as England Finder coordinates. The resulting slides
138 are deposited in the Microfossil Collection of the Universidad Nacional de Salta,
139 Argentina, under the acronym CNS-M.

140

141 **4. Results and discussion**

142 Well-preserved palynomorphs were obtained from a single palynological sample
143 recovered from the Yacoraite Formation outcrops. Fifty eight species have been
144 identified, 38 angiosperms (two monocot families, eleven eudicot families and several
145 of uncertain affinity), 7 gymnosperms of two families, and within the spores, 2
146 Marchantiophyta, 6 ferns of 6 families and 4 of uncertain affinity, along with 1
147 freshwater algae). Brief description and comments of taxa are listed at Appendix 1. The
148 identified palynomorphs and their botanical affinities are listed in Table 1, and
149 illustrated in figures 3 through 5.

150 Based on a quantitative analysis with an N = 306 specimens, the palynological
151 assemblage as a whole is dominated by angiosperm pollen grains (52.4%), followed by
152 Monilophyta (24.5%), gymnosperm pollen grains (11.5%), and freshwater algae (4.8%),
153 while bryophytes and fungi are scarce with 2.5 and 2.9%, respectively (see Appendix
154 2).

155 The sample is devoid of marine palynomorphs.

156 **4.1. Comparisons**

157 **4.1.1. Argentina**

158 Within the Salta Group, as previously mentioned, Moroni (1982) reported a palynoflora
159 for the Yacoraite Formation, recovered from Lomas de Olmedo sub-basin sediments.
160 Two different assemblages were identified. One with the presence of *Aquilapollenites*
161 *magnus*, *Crassitriapertites brasiliensis* (= *Crassitricolporites brasiliensis*), *Gabonisporis*
162 *vigorousii*, *Psilastephanosporites brasiliensis*, the genus *Zlvisporis*, appearing together
163 with numerous poliplicate grains (i.e. *Ephedripites* complex), and in less amounts
164 *Tricolpites* sp., *Ulmoideipites* sp. (= *Verrustephanoporites*), algae and dinoflagellate
165 cysts, to which Moroni (1982) assigned a Maastrichtian age. The other one, was
166 described as a pollen assemblage with high Ulmaceae values, tricolpate, tricolporate,
167 monocolpate, and triporate, along with *Pediastrum* sp. and *Mtchedlishvilia saltenia*
168 Moroni 1984, and assigned to the Danian (Moroni, 1982, 1984). Our pollen assemblage
169 shares all the characteristics and taxa described by Moroni (1982) for the Maastrichtian
170 assemblage.

171 Quattrocchio et al. (2005) reported a palynoflora from transitional levels between the
172 Los Blanquitos and Lecho formations (top of the Pirgua subgroup and base of the
173 Balbuena Subgroup, respectively). This assemblage, dated as Campanian to early
174 Maastrichtian, is characterized by typical Mesozoic taxa like *Classopollis* spp.,

175 *Callialasporites* spp. and *Todisporites major* together with angiosperm taxa previously
176 reported in younger units from tropical South America and Africa such as
177 *Reitrescolpites saturum*, *Retistephanocolpites regularis* and *Spinitricolpites*
178 *jennerclarkei*. The last two taxa, together with *Verrustephanoporites simplex* (all typical
179 of younger floras), are also present in our assemblage. The lack of the typical Mesozoic
180 elements mentioned by Quattrocchio et al. (2005) would probably indicate that the
181 Yacoraite flora is younger than that Campanian–early Maastrichtian flora. A
182 preliminary conclusion subject to the fact that more samples are needed to make
183 stronger biostratigraphic inferences based on the palynofloral diversity.
184 On the other hand, palynological assemblages from the Tunal Formation (Danian),
185 which overlies the Yacoraite Formation, were studied by Quattrocchio and Volkheimer
186 (1988), Quattrocchio et al. (1988) and Volkheimer et al. (2006) at southern Salta
187 Province. These palynofloras are referred to the Danian *Mtchedlishvilia saltenia*
188 palynozone (Quattrocchio et al., 2000), with *M. saltenia*, associated with *Pandaniidites*
189 *texus*, *Gemmatricolpites subsphaericus* and *Clavatricolpites* sp. cf. *C. gracilis*.
190 *Verrustephanoporites simplex* dominates (62%–88.5 %) the Tunal Formation
191 palynological assemblages (Quattrocchio et al., 1988; Volkheimer et al., 2006). The
192 species shared with the Tunal Formation are *Apiculatisporis* sp. cf. *A. charahuillaensis*,
193 *Gabonisoris vigorouxii*, and *Polypodiaceoisporites* sp. cf. *P. retirugatus*, and the
194 genera *Gleicheniidites* and *Zlivisporis*, among the spores; also some species from the
195 *Ephedripites* complex and the genus *Podocarpidites*, among the gymnosperms; and
196 *Gemmatricolpites subsphaericus*, *Tricolpites reticulatus* and *Verrustephanoporites*
197 *simplex*, together with the genera *Myriophyllumpollenites*, *Retitricolporites*, *Rousea* and
198 *Tricolpites* among the angiosperms. However, in our assemblage the typical Danian

199 species are absent, and percentage of *Verrustephanoporites simplex* in the palynoflora
200 we describe is far less (13.2 %) than that recorded for the Tunal Formation.
201 Finally, a palynoflora from the Ciénaga del Río Huaco Formation in the San Juan
202 Province (central Argentina) was briefly described (Maastrichtian; Limarino et al.,
203 2000). Fourteen species were identified, of which only *Gabosisporites vigorouxii* and
204 *Ephedripites jansonii* are shared with our palynoflora.

205 **4.1.2. Bolivia**

206 Pérez Leytón (1987) reported palynological assemblages from two Upper Cretaceous
207 formations from southwestern Bolivia. The assemblage from the Chaunaca Formation
208 (Santonian–early Campanian) is characterized by several species from the *Ephedripites*
209 complex, monocolpate, tricolpate and tricolporate grains and the presence of
210 *Normapolles* spp. Our assemblage shares with the Chaunaca Formation palynoflora the
211 presence of the genera *Podocarpidites*, *Aquillapollenites*, and a diverse *Ephedripites*
212 complex. On the other hand, the palynoflora from the El Molino Formation
213 (Maastrichtian) was described based on a single sample with 16 species of ferns and
214 bryophyte spores, gymnosperm and angiosperm pollen grains, all in relatively equal
215 proportions. The taxon *Gabonispuris vigorouxii* and the genus *Rhoipites* are considered
216 abundant (appearing in more than 10 % of the total assemblage). In our assemblage the
217 angiosperms are largely dominant, the genus *Rhoipites* is absent, and *Gabonispuris*
218 *vigorouxii* has a lower proportion (1.8%).

219 Another palynological work from Bolivia is that of Vajda-Santivanez (1999) who
220 studied the Eslabón (Maastrichtian) and Flora (Paleocene) formations from
221 northwestern Bolivia. Both assemblages are of continental origin with their composition
222 resembling that of the Late Cretaceous Palmae Province of Herengreen et al. (1996). A
223 lacustrine paleoenvironment of deposition is suggested and a warm and humid (tropical)

224 climate (Vajda-Santivanez, 1999). The Yacoraite palynoflora is very similar to that of
225 the Maastrichtian Eslabón Formation in terms of proportions of each group, with
226 angiosperms being clearly dominant, and with equivalent amounts of gymnosperms and
227 ferns. Also shares the presence of several taxa, such as *Gabonisporis vigorouxii*, a
228 diverse *Ephedripites* complex, together with the angiosperm species *Aquillapollenites*
229 *magnus*, *Crassitricolporites brasiliensis*, *Ulmoideipites krempii* (= *Verrustephanoporites*
230 *simplex*), and the genera *Retitricolporites*, *Tricolpites*, *Tricolporites*. The latter two
231 genera and *Ulmoideipites krempii* (= *Verrustephanoporites simplex*) are also shared with
232 the Flora Formation.

233 **4.1.3. Brazil**

234 Regali et al. (1974a, b) carried out a detailed palynological study from different
235 Brazilian Cretaceous–Neogene coastal basins. Within them, those located further south
236 would have developed in mid-latitudes (Campos Basin and Espírito Santo Basin). The
237 authors described the *Crassitricolporites brasiliensis* superzone for the mid Campanian–
238 Maastrichtian. The palynoflora of the Yacoraite Formation presents many taxa shared
239 with that palynozone. The species *Crassitricolporites brasiliensis*, *Aquillapollenites*
240 *magnus*, *Ulmoideipites krempii* (= *Verrustephanoporites simplex*), *Ephedripites*
241 complex (e.g., *Gnetaceaepollenites*, *Ephedripites*), *Tricolpites* spp., the genera
242 *Zlivisporis*, *Polypodiaceoisporites* (*P. potoniaei*) and *Triorites* (*T. africaensis*),
243 *Psilastephanoporites brasiliensis*, *Gabonisporis vigorouxii* and *Crassitricolporites*
244 *brasiliensis* are shared.

245

246 **4.2. Biostratigraphy and age determination**

247 Some of the recorded species provide reliable information to constrain the age of the
248 palynoflora studied here. *Aquillapollenites magnus*, *Gabonisporis vigorouxii* and

249 *Crassitricolporites brasiliensis* are typically recorded in Maastrichtian deposits from
250 Brazil (Vajda-Santivanez, 1999), where they usually go extinct at the K–Pg boundary
251 (Vajda and Bercovici, 2012). However, *Gabonisporis vigorouxii* is recorded in
252 Santonian through middle Eocene deposits from Argentina (Quattrocchio, 1978;
253 Moroni, 1982; Limarino et al., 2000; Volkheimer et al., 2006, 2007; Vallati, 2010). On
254 the other hand, *Gemmatricolpites subsphaericus* presents its earliest records in Danian
255 deposits of Argentina (Archangelsky, 1973; Quattrocchio and Volkheimer, 1988;
256 Volkheimer et al., 2006, 2007). *Retistephanocolpites regularis* and *Spinetricolpites*
257 *jennerclarkei* have been previously recorded in Campanian to early Maastrichtian
258 sedimentary rocks from the Salta Basin (Quattrocchio et al., 2005), and the latter was
259 also found in the Danian Olmedo Formation (Narváez, 2009). Finally, *Syncolporites*
260 *marginatus* originally described for the Maastrichtian of Nigeria (Van Hoeken-
261 Klinkenberg, 1964), is recorded from Campanian to Paleogene deposits (e.g., Schrank,
262 1984; Jaramillo and Dilcher, 2001; Eisawi et al., 2012; Cole et al., 2017).
263 Quattrocchio et al. (2000) proposed a palynological zonation for Paleocene–Eocene
264 strata from the Salta Basin. The Danian *Mtchedlishvilia saltenia* Zone (Quattrocchio et
265 al., 2000) is defined by the first appearance of *M. saltenia*, associated with
266 *Pandaniidites texus*, *Gemmatricolpites subsphaericus* and *Clavatricolpites* sp. cf. *C.*
267 *gracilis*. In the palynoflora studied here, *M. saltenia* is absent, and only *G.*
268 *subsphaericus* was identified. On the other hand, Regali et al. (1974a, b) described the
269 *Crassitricolporites brasiliensis* superzone and the distribution of many taxa that are also
270 present in the flora of the Yacoraita Formation: *C. brasiliensis* and *Aquilapollenites*
271 *magnus* are present since the early Maastrichtian and *Verrustephanoporites simplex* (a
272 quite abundant -13.2%- element in our assemblage) since the late Maastrichtian through
273 the Oligocene. Species from the *Ephedripites* complex (e.g., *Gnetaceaepollenites*,

274 *Ephedripites*) and *Tricolpites* spp. appear in Upper Cretaceous levels. Moreover, the
275 genera *Zlivisporis*, *Polypodiaceoisorites* (*P. potoniaei*; Campanian through Miocene
276 from Brazil) and *Triorites* (*T. africaensis*; recorded in Cenomanian deposits), and the
277 species *Psilastephanoporites brasiliensis* (Coniacian to Maastrichtian), *Gabonisoris*
278 *vigorousii* (Santonian to Maastrichtian) and *Crassitricolporites brasiliensis* (present in
279 Campanian–Maastrichtian strata). Hence, based on the distribution and relative
280 abundance of the mentioned taxa, a late Maastrichtian age could be suggested for the
281 palynological assemblage studied herein, although more samples are needed to confirm
282 this.

283 **4.3. Paleoenvironment and paleoclimate**

284 Several works discuss the depositional environment of the Yacoraite Formation (e.g.,
285 Marquillas, 1985, 1986; Salfity and Marquillas, 1994; Marquillas et al., 2005). Whole-
286 rock sample analyses from the Metán sub-basin show carbon and oxygen isotope
287 behavior typical of marine to freshwater transition within this formation (Marquillas et
288 al., 2007). At the sampling area, Zelaya et al. (2014a, b) identified four facies
289 associations that can be interpreted as a whole, as a paleoenvironment with shallow
290 conditions near coastline settings, alternating with sub-littoral deposits. The
291 palynological sample studied herein corresponds to F8 (dark gray to black siltstones and
292 claystones), included in FA2 interpreted as deposits of tidal channels associated with
293 intertidal mud flats, grading into a restricted sedimentary environment (Zelaya et al.,
294 2014a, b).

295 Shallow marine conditions and localized mixing of fresh and brackish water are
296 indicated by the invertebrate fossil record (e.g. euryhaline association composed of
297 ostracods, bivalves, gastropods, and foraminifera and different types of charophytes;
298 Marquillas et al., 2005); which is consistent with the palynological evidence of Papú

299 and Melendi (1984) who reported the presence of the aquatic fern *Azolla cretacea*
300 together with dinoflagellate cysts in the unit. However, some authors still question the
301 presence of a marine environment, or restrict it to some parts of the basin. For example,
302 Musacchio (1972) reported the presence of non-marine ostracod and charophytes at the
303 Tres Cruces sub-basin. Similarly, based on the ichthyofaunal records, Cione and Pereira
304 (1985) determined that the absence of shark remains denote that the water bodies
305 present in this unit would not correspond to a marine environment. Furthermore, Starck
306 (2011) considers that the fossil evidence (e.g., foraminifera, mollusks, fishes) consists
307 mainly of non-diagnostic elements compatible with brackish waters, and that the most
308 likely paleoenvironmental scenario would be that of a distant marine connection
309 through the Bolivian Andean Basin. A connection that would have varied from
310 moments of greater connection to others of total isolation during the deposition of the
311 Yacoraite Formation (Starck, 2011).

312 De Valais and Cónsole-Gonella (2018) studied two ichnofossiliferous localities of the
313 Yacoraite Formation, Maimará (Tres Cruces sub-basin) and Quebrada del Tapón
314 (Alemanía sub-basin), the latter close to the Don Otto mine. Based on the record of
315 avian footprints, they determined that in the Tres Cruces sub-basin, the facies
316 correspond to a coastal environment, while in Quebrada del Tapón to a continental
317 environment. This is consistent with our results, which show a distinctly continental
318 palynological assemblage, without marine palynomorphs. Nevertheless, further studies
319 with more palynological samples would help to make more conclusive
320 paleoenvironmental inferences.

321 Considering the presence of *Verrustephanoporites simplex* in the described palynoflora,
322 a paleoenvironment similar to the modern Transitional Forest (Cabrera, 1976) or
323 Piedmont Semi-deciduous Rain Forest (Oyarzabal et al., 2018) within the Yungas

324 Province (350–500 m a.s.l.) could be preliminary inferred. *V. simplex* is related to
325 *Phyllostylon rhamnoides* (Ulmaceae), a dominant tree in the Transitional Forest, as it
326 was recognized by Quattrocchio et al. (1988) for the time of deposition of the Danian
327 Tunal Formation in the Salta Basin. The climate of the Yungas Province is humid and
328 warm, with summer rainfalls, some frost in winter and annual precipitations of
329 approximately 700–1000 mm (Cabrera, 1976; Oyarzabal et al., 2018).

330 The presence of Haloragaceae (*Myriophyllumpollenites* spp.) together with
331 Marsileaceae (*Gabonisporis vigorouxii*) and scarce *Ovoidites* sp. indicate a lacustrine
332 environment during the deposition of at least part of the Yacoraite Formation.

333 Moreover, the presence of *Gemmatricolpites subsphaericus* (Aquifoliaceae),
334 *Retistephanocolpites regularis*, *Spinitricolpites jennerclarkei* (Valerianaceae?-
335 Verbenaceae) and pollen grains related to Arecaceae (*Arecipites*, *Spinozonocolpites*),
336 suggest warm and at least locally humid conditions.

337 The lack of gymnosperm taxa such as *Classopollis* and *Callialasporites* that were
338 present in the transitional levels between the Los Blanquitos and Lecho formations
339 (Campanian–early Maastrichtian; Quattrocchio et al., 2005) would suggest less dry
340 conditions for the late Maastrichtian but as abovementioned, more samples are needed
341 to confirm this.

342 **4.4. Paleobiogeography**

343 Herngreen et al. (1996) established two Late Cretaceous palynofloral provinces in South
344 America: *Palmae* and *Proteacidites/Nothofagidites* provinces at low and high latitudes,
345 respectively. The increase and diversity of monocolpate *Palmae*-type pollen and
346 appearance of proteacean pollen, typifies the *Palmae* Province, together with a low
347 abundance of gymnosperm pollen (e.g., decrease of *Ephedripites*-type species), and
348 particularly, the absence of bisaccate pollen (Herngreen and Chlonova, 1981; Vajda-

349 Santivanez, 1999). On the other hand, the southern *Proteacidites/Nothofagidites*
350 Province is defined by the presence of nothofagaceous pollen, together with
351 *Beaupreaidites*, *Proteacidites* and *Peninsulapollis* taxa (with proteaceous affinity), and
352 the abundance of podocarpaceous pollen (Herngreen et al., 1996). Later, Vajda and
353 Bercovici (2012, 2014) defined a new zone between the two mentioned provinces, the
354 transitional or mixed flora, with a mixed composition but also containing some endemic
355 taxa.

356 The Yacoraite palynoflora developed at mid-latitudes, and is located on the northern
357 edge of the transitional zone consisting of mixed floras (Vajda and Bercovici, 2012,
358 2014). The palynological assemblage has some characteristics that link it to the Palmae
359 Province, such as the presence of *Gabonispores vigorouxii*, *Aquillapollenites magnus*,
360 and *Crassitricolporites brasiliensis*, together with the low abundance of gymnosperm
361 pollen. But also has other characteristics that exclude it from that phytogeographic unit,
362 such as the less diverse record of palmae-pollen (*Arecipites* sp. cf. *A. tenuixinus*, and
363 *Spinizonocolpites* sp.), and the scarcity of proteaceous pollen. As for bisaccate pollen-
364 producing conifers, *Podocarpidites* sp. is very scarce (0.8%). The bond of the Yacoraite
365 flora with the *Proteacidites/Nothofagidites* Province is even less clear due to the
366 scarcity of Proteaceae (0.7%) and *Nothofagidites* sp. (1.9%). Representatives of the
367 Nothofagaceae have anemophylous pollination, and its pollen can be dispersed over
368 great distances, even thousands kilometers from its area of origin. This was recorded by
369 Muñoz et al. (2017) who found *Nothofagus* pollen grains in Entre Ríos Province
370 (central-eastern Argentina), which is at least 1500 km away from the Andean
371 Patagonian forests where this taxon currently lives, being considered an extra-regional
372 element of that pollen spectrum.

373 Moreover, several taxa present in the Yacoraite palynoflora such as *Crassitricolporites*
374 *brasiliensis*, *Aquillapollenites magnus*, *Gabonisporis vigorouxii* and
375 *Verrustephanoporites simplex* (= *Ulmoideipites krempii*), are also found in Maastrichtian
376 deposits from northern South America (Brazil), included by Regali et al. (1974a, b) in
377 the *Crassitricolporites brasiliensis* palynozone.

378

379 **5. Concluding remarks**

380 A single palynological sample from the Yacoraite Formation presented here yields
381 palynomorphs of relatively good preservation that provide relevant information for a
382 unit where the palynological records are scarce compared to other Maastrichtian
383 formations in Argentina (mainly from high latitudes in Patagonia). This palynological
384 record infers a late Maastrichtian age for the unit in the outcrops of the Valle del Tonco.
385 The palynological assemblage does not present a clear relationship with any of the
386 phytogeographic provinces described for the Late Cretaceous. Although it presents
387 more similarities with the Palmae Province, it better fits with the transitional or mixed
388 flora. The study of palynofloras from mid-latitudes can help understand the transition
389 between tropical and temperate floras in South America during the Late Cretaceous.

390

391 **Acknowledgements**

392 This research was partially funded by the Consejo Nacional de Investigaciones
393 Científicas y Técnicas (Project PIP CONICET 11220200101738CO).

394 We want to express our gratitude to Regional Noroeste of Gerencia Exploración de
395 Materias Primas, Comisión Nacional de Energía Atómica (CNEA) for the palynological
396 samples cession and for giving permission to publish these results; especially to Eng.
397 Inés Torino Aráoz and Geol. Luis Ferreira for their attentive collaboration. The editor

398 Eduardo Koutsoukos and two anonymous reviewers are acknowledged for providing
399 constructive comments that improved the manuscript.

400

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644

645



646 **Captions**

647

648 **Table 1.** Taxonomic and botanical affinity of the taxa recorded in the Yacoraite
649 Formation at the Don Otto mine vicinity.

650

651 **Fig. 1.** Location map of the study area. A. General location, B. Salta Province, C.
652 Location of study área, D. Geological map of the Valle del Tonco (modified from
653 Gorustovich et al., 2013).

654

655 **Fig. 2.** Lithostratigraphic section with palynological sample location (modified from
656 Zelaya et al., 2014a).

657

658 **Fig. 3.** Spores and gymnosperm pollen grains of the Yacoraite Formation. A.
659 *Apiculatisporis* sp. cf. *A. charahuillaensis*, CNS-M 001/C: M43/3; B. *Baculatisporites*
660 *comaumensis*, CNS-M 001/G: G39/0; C. *Concavissimisporites* sp., CNS-M 001/B:
661 Z30/1; D. *Fisciniasporites* sp., CNS-M 001/B: X57/4; E. *Gabonisoris vigorouxii*,
662 CNS-M 001/E: B45/2; F. *Gleicheniidites* sp., CNS-M 001/G: U58/1; G.
663 *Microreticulatisporites* sp., CNS-M 001/G: Y36/0; H, *Polypodiaceoisporites*
664 *retirugatus*, CNS-M 001/E: A24/3; I. *Verrucosisporites* sp. 1, CNS-M 001/G: G39/2;
665 J. *Verrucosisporites* sp. 2, CNS-M 001/A: S45/0; K. *Zlivisporis reticulatus*, CNS-M
666 001/D : A59/4; L. *Zlivisporis* sp., CNS-M 001/A: L28/4; M. *Equisetosporites* sp. cf. *E.*
667 *concinus*, CNS-M 001/G: P31/4; N. *Equisetosporites* sp. cf. *E. notensis*, CNS-M
668 001/G: Y59/1; O. *Gnetaceaepollenites jansonii*, CNS-M 001/C: T61/2; P.
669 *Gnetaceaepollenites* sp. cf. *G. barghoornii*, CNS-M 001/G: N43/2; Q.
670 *Gnetaceaepollenites* sp., CNS-M 001/F: G51/4; R. *Singhia multicostata*, CNS-M

671 001/H: G51/1; S. *Podocarpidites* sp., CNS-M 001/B: Z28/0; T. *Ovoidites* sp., CNS-M
672 001/B: Y29/1. Scale bar: 10µm.

673

674 **Fig. 4.** Angiosperm pollen grains of the Yacoraite Formation. A–B. *Aquillapollenites*
675 *magnus*; A. CNS-M 001/B: Z34/1; B. CNS-M 001/G: L58/1; C–D, *Crassitricolporites*
676 *brasiliensis*; C. CNS-M 001/E: J46/1; D. SEM; E. *Cupanieidites* sp. cf. *C. orthoteichus*,
677 CNS-M 001/F: L40/3; F. *Gemmatricolpites subsphaericus*, CNS-M 001/H: S51/0; G–H.
678 *Gemmatricolpites pergemmatus*; G. CNS-M 001/H: T48/3; H, SEM; I. *Margocolporites*
679 sp. cf. *M. lihokus*, CNS-M 001/E: K47/0; J. *Diporites* sp., CNS-M 001/A: Z26/3; K.
680 *Myriophyllumpollenites* sp. in Quattrocchio and Volkheimer 1988, CNS-M 001/E:
681 U44/0; L. *Penetetrapites* sp. cf. *P. mollis*, CNS-M 001/E: Z20/4; M. *Peninsulapollis*
682 *gillii*, CNS-M 001/G: Y34/3; N. ?*Psilabrevitricolpites* sp., CNS-M 001/G: Z26/2; O.
683 *Psilastephanoporites brasiliensis*, CNS-M 001/E: G55/4; P. *Psilatricolpites* sp., CNS-M
684 001/E: S37/2; Q. *Arecipites* sp. cf. *A. tenuixinus*, CNS-M 001/H: W33/0; R.
685 *Liliacidites variegatus*, CNS-M 001/G: Z40/4; S. *Psilatricolporites* sp. 1, CNS-M
686 001/E: F50/2; T. *Psilatricolporites* sp. 2, CNS-M 001/H: D50/0; U.
687 *Retistephanocolpites regularis*, CNS-M 001/H: T27/3; V. *Nothofagidites* sp., CNS-M
688 001/H: T53/2; W–X. *Retitricolporites* sp. cf. *R. multibaculates*; W, CNS-M 001/A:
689 W28/0; M43/3; X, CNS-M 001/B: H30/1; Y. *Retitricolporites* sp., CNS-M 001/F:
690 O57/0. Scale bar: 10µm.

691

692 **Fig. 5.** Angiosperm pollen grains of the Yacoraite Formation. A. *Rousea patagonica*,
693 CNS-M 001/E: H49/3; B. *Spinitricolpites jennerclarkei*, CNS-M 001/D: E34/1; C.
694 *Spinizonocolpites* sp., CNS-M 001/G: C28/3; D. *Syncolporites* sp. cf. *S. marginatus*,
695 CNS-M 001/H: J51/3; E. *Syncolporites* sp., CNS-M 001/H: Q40/3; F–G.

696 *Striatricolporites* sp.; F, CNS-M 001/E: Y39/0; G, CNS-M 001/E: L32/2; H–I.
697 *Tricolpites* sp. cf. *T. bathyreticulatus*; H, CNS-M 001/A: N49/3; I, SEM; J. *Tricolpites*
698 *reticulatus*, CNS-M 001/G: S47/2; K. *Tricolpites* sp. cf. *T. cellatus*, CNS-M 001/A:
699 V47/0; L. *Tricolpites* sp., CNS-M 001/H: G35/0; M–N. *Verrustephanoporites simplex*;
700 M, CNS-M 001/E: R51/0; N, CNS-M 001/E: B21/4; O. *Triorites* sp. cf. *T. tenuixinus*,
701 CNS-M 001/E: C36/4; P. *Verrutricolpites* sp., CNS-M 001/C: Q40/0; Q. *Tricolporites*
702 sp. 1, CNS-M 001/H: U44/3; R. *Tricolporites* sp. 2, CNS-M 001/H: B20/0; S.
703 *Tricolporites* sp. 3, CNS-M 001/B: M36/4; T. *Tricolporites* sp. 4, CNS-M 001/E:
704 C19/0. Scale bar: 10 μ m.



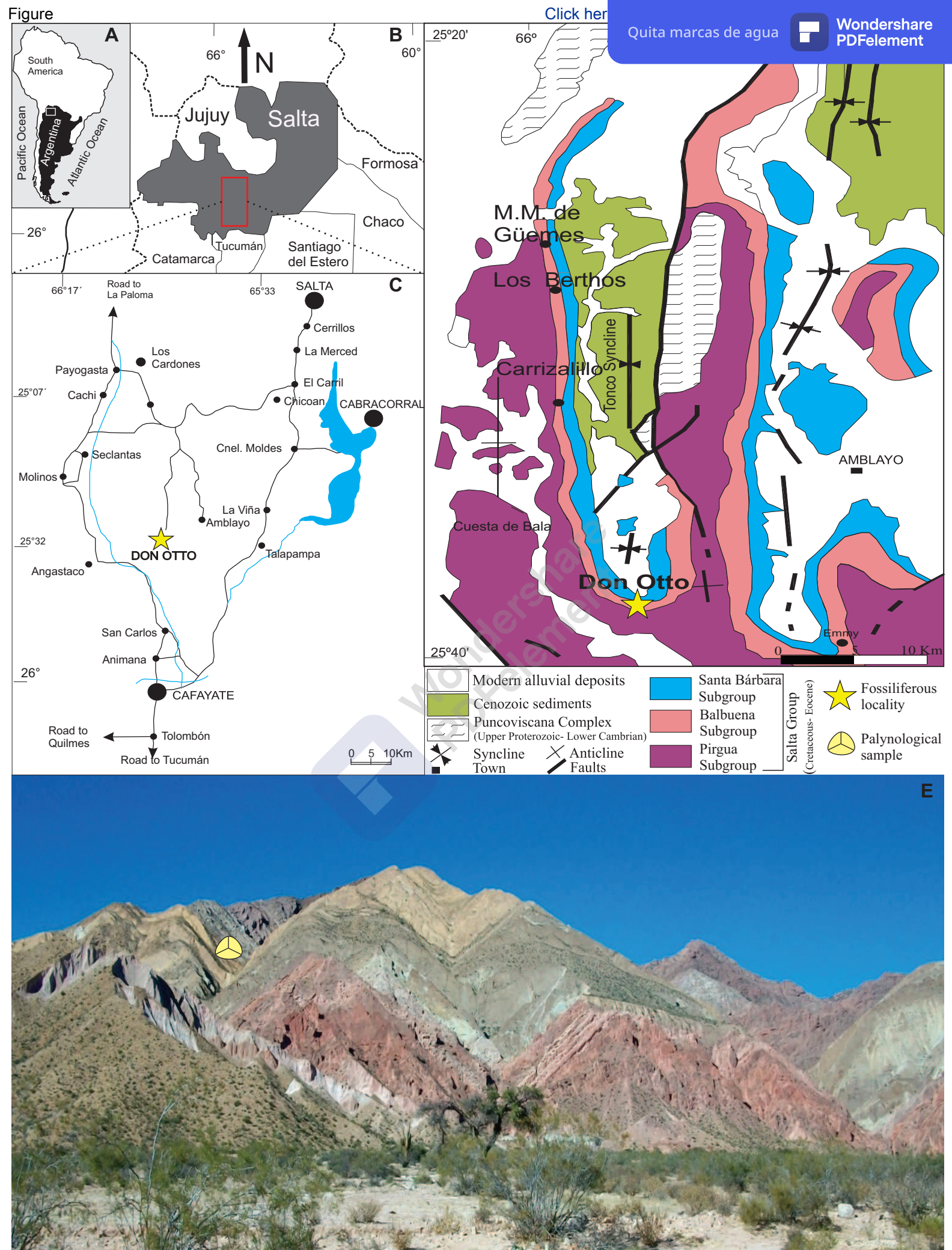
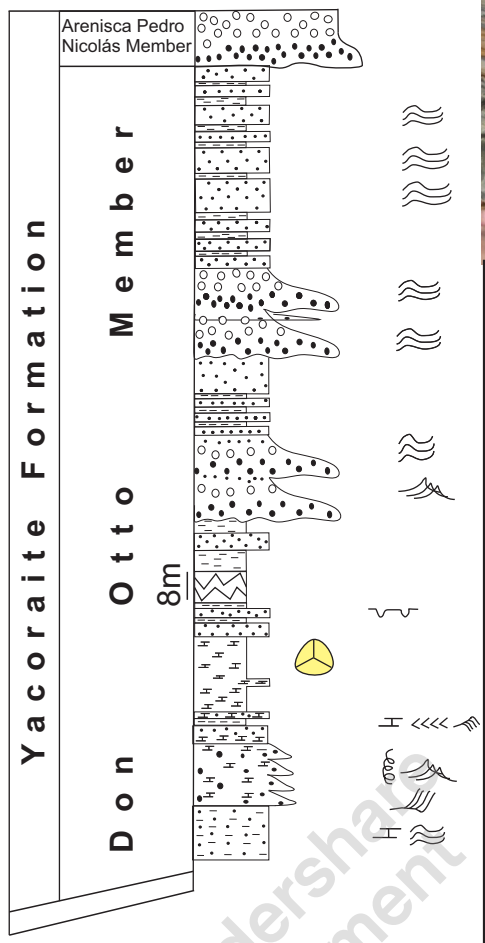
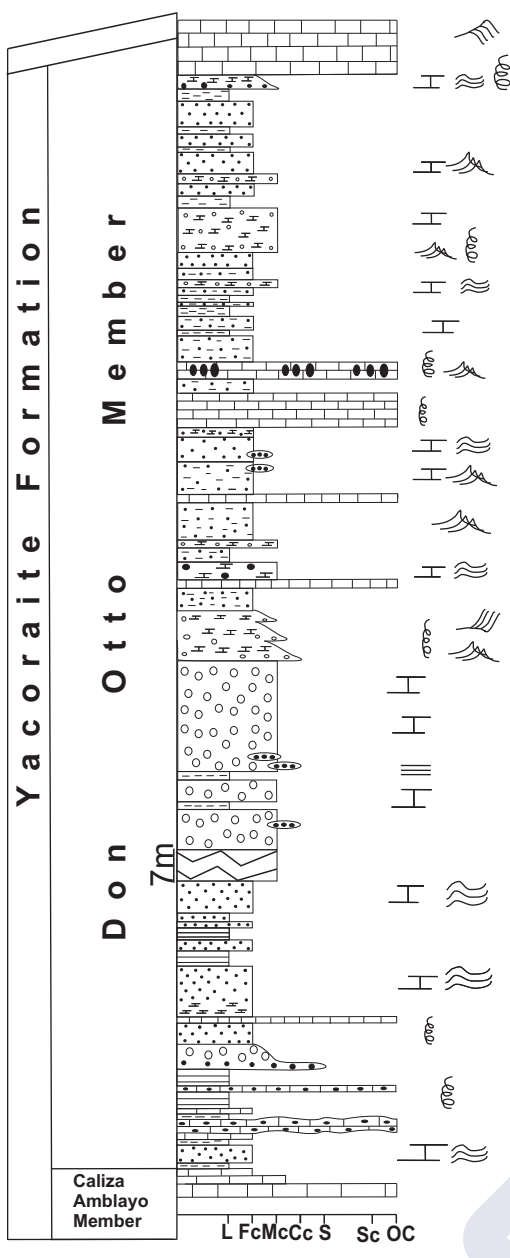


Figure 1. Location of the study area



	Limestone
	Fine Clay
	Medium Clay
	Coarse Clay
	Sandstones
	Oolitic-calcite
	Sandy-calcite
	Inter-stratified sandstone and limestones
	Flat lamination
	Wavy lamination
	Criss-cross lamination
	Tangential lamination
	Herringbone
	Mud cracks
	Ripple marks
	Bioturbation
	Calcareous layers
	Sandstone lenses
	Palynological sample

Wondershare PDFelement

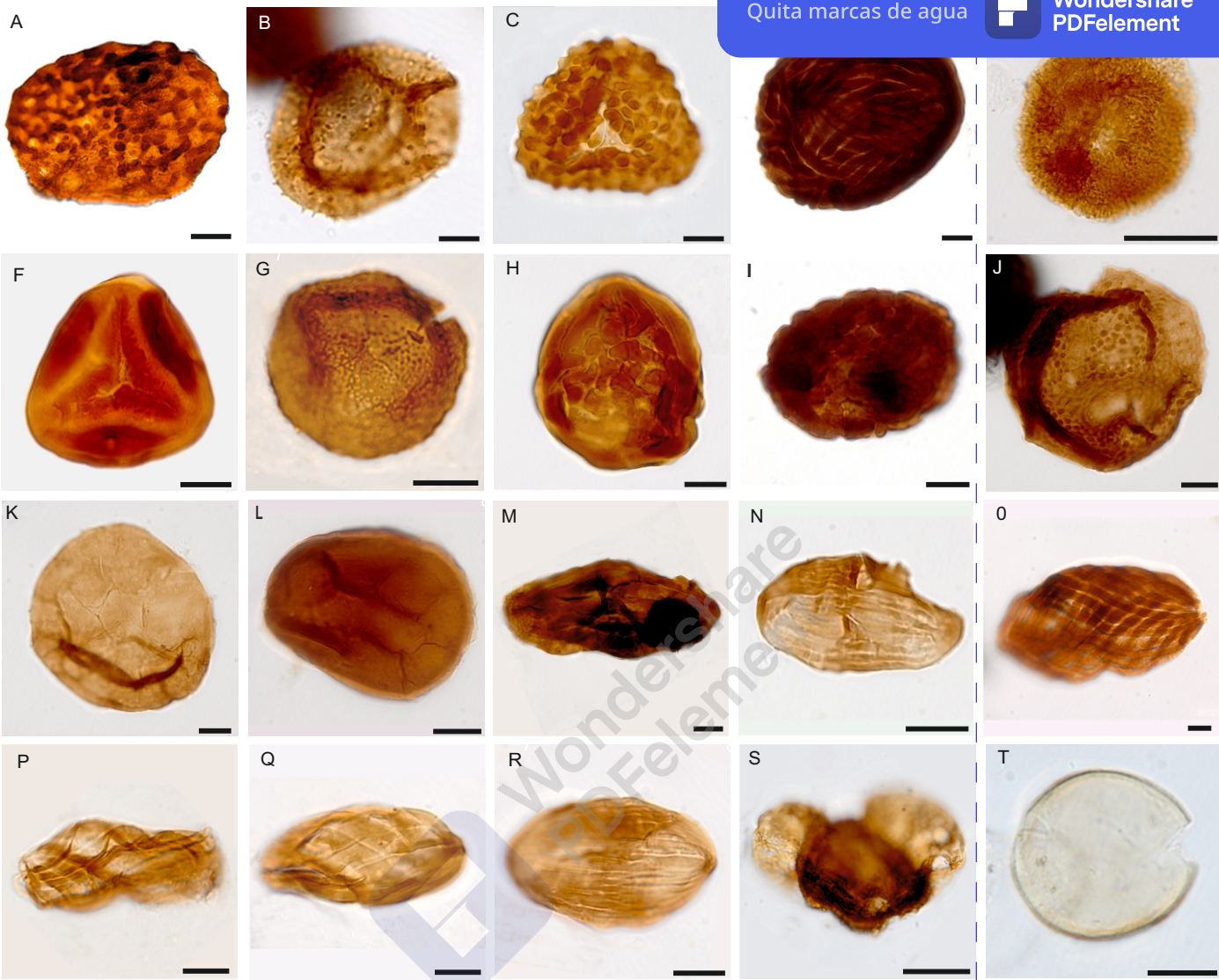
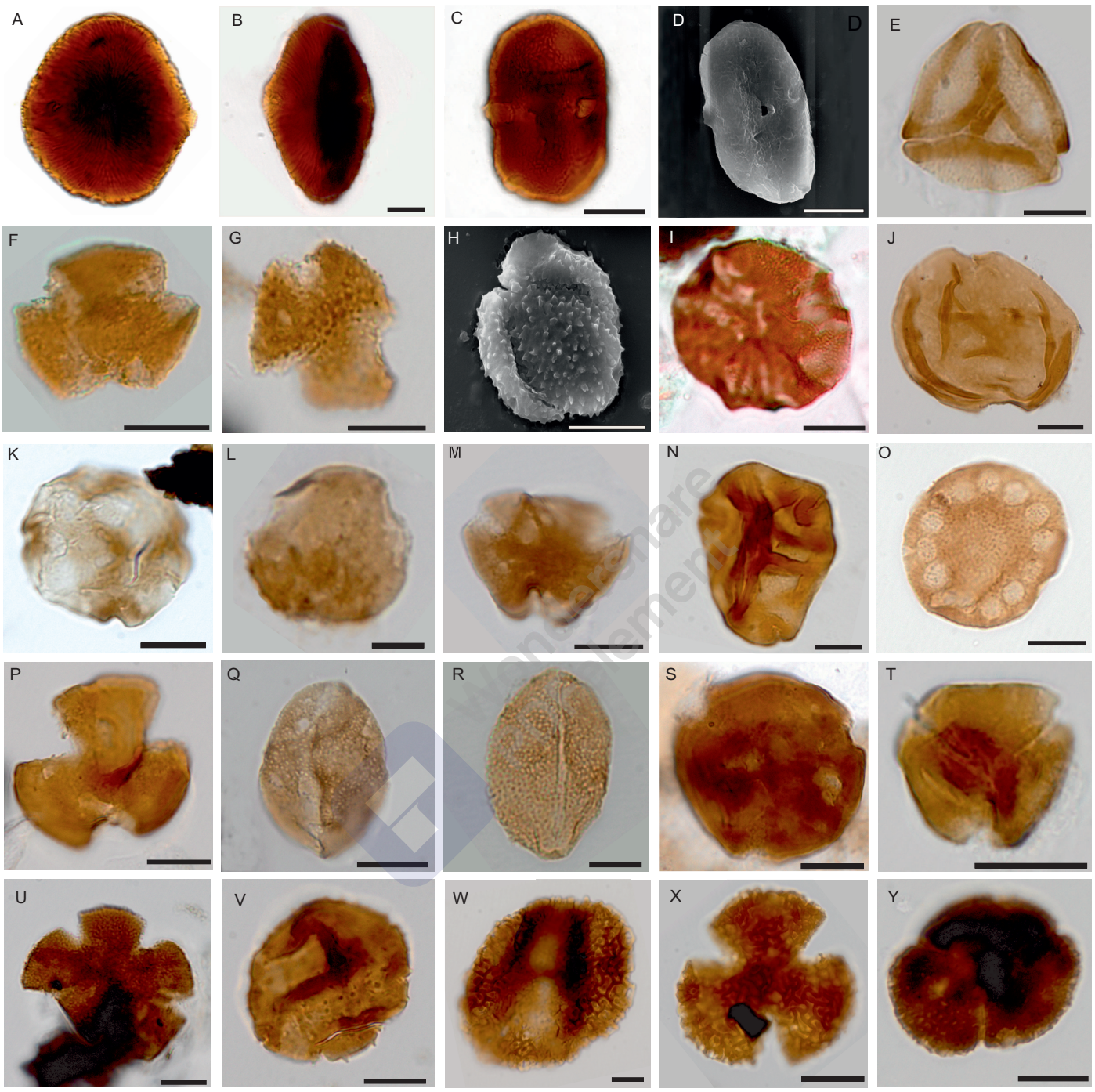


Figure 3



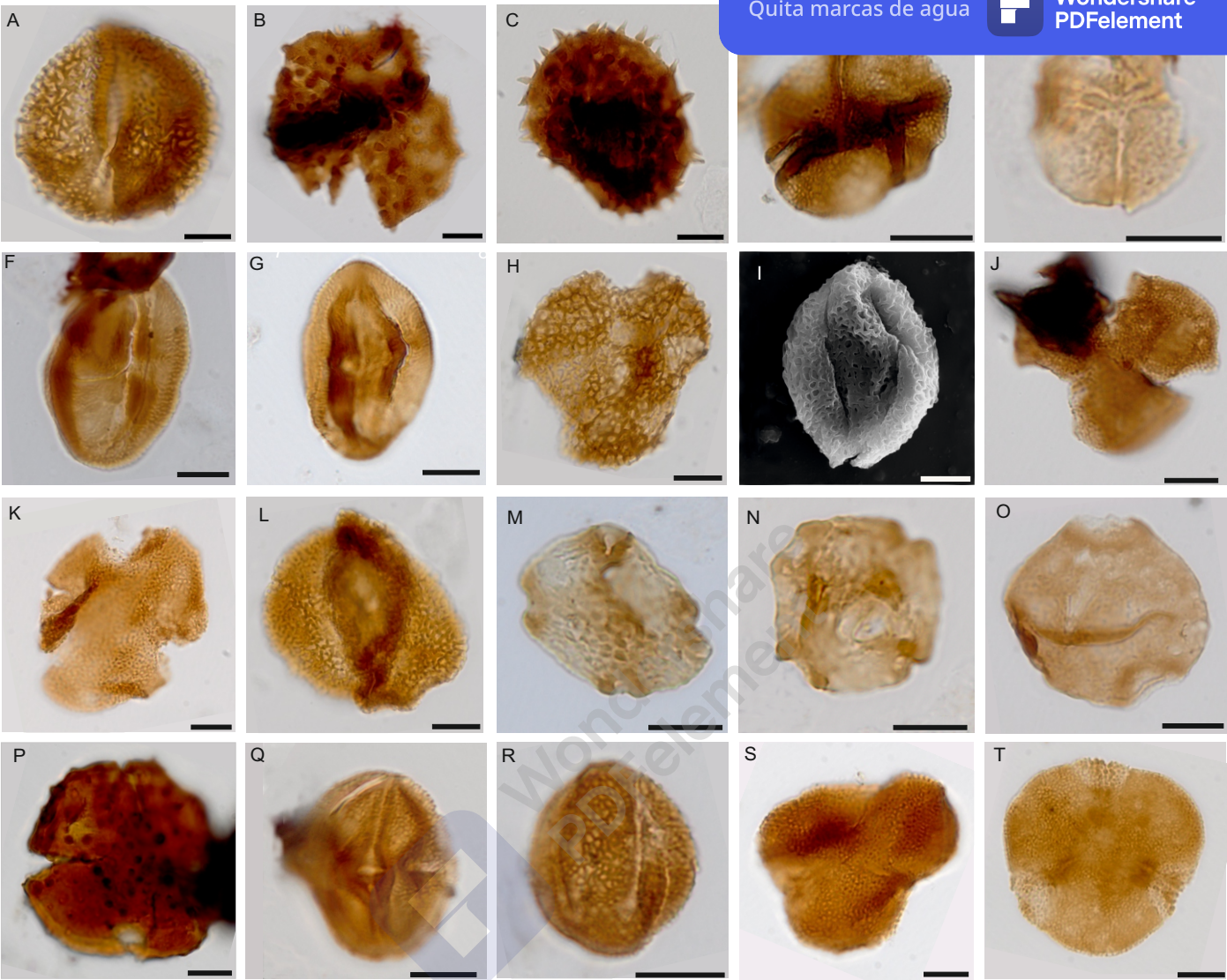


Figure 5



Group	Family	Species	
Spores			
Marchantiophyta		<i>Zlivisporis reticulatus</i> (Pocock) Pacltova and Simoncsics 1970 (Fig. 3K)	
		<i>Zlivisporis</i> sp. (Fig. 3L)	
Monilophyta	Schizaeaceae	<i>Fisciniasporites</i> sp. (Fig. 3D)	
	Polypodiaceae	<i>Polypodiaceoisporites retirugatus</i> Muller 1968 (Fig. 3H)	
	Osmundaceae	<i>Baculatisporites comaumensis</i> (Cookson 1953) Potonie 1956 (Fig. 3B)	
	Gleicheniaceae	<i>Gleicheniidites</i> sp. (Fig. 3F)	
	Dicksoniaceae	<i>Concavissimisporites</i> sp. (Fig. 3C)	
	Marsileaceae	<i>Gabonisoris vigorouxii</i> Boltenhagen 1967 (Fig. 3E)	
Incertae sedis		<i>Microreticulatisporites</i> sp. (Fig. 3G)	
		<i>Verrucosisporites</i> sp. 1 (Fig. 3I)	
		<i>Verrucosisporites</i> sp. 2 (Fig. 3J)	
		<i>Apiculatisporis</i> sp. cf. <i>A. charahuillaensis</i> (Fig. 3A)	
Algae	Zignemataceae	<i>Ovoidites</i> sp. (Fig. 3T)	
Fungi		various spores	
Gymnosperm pollen grains			
Coniferophyta	Podocarpaceae	<i>Podocarpidites</i> sp. (Fig. 3S)	
Gnetophyta	Ephedraceae	<i>Equisetosporites</i> sp. cf. <i>E. concinnus</i> Singh 1964 (Fig. 3M)	
		<i>Equisetosporites</i> sp. cf. <i>E. notensis</i> (Cookson) Romero 1977 (Fig. 3N)	
		<i>Gnetaceapollenites jansonii</i> (Pocock) Lima 1980 (Fig. 3O)	
		<i>Gnetaceapollenites</i> sp. cf. <i>G. barghoornii</i> (Pocock) Lima 1980 (Fig. 3P)	
		<i>Gnetaceapollenites</i> sp. (Fig. 3Q)	
		<i>Singhia multicostata</i> (Brenner) Lima 1980 (Fig. 3R)	
Angiosperm pollen grains			
Monocots	Arecaceae	<i>Arecipites</i> sp. cf. <i>A. tenuixinus</i> Leffingwell 1971 (Fig. 4Q)	
	Liliaceae	<i>Spinizonocolpites</i> sp. (Fig. 5C)	
Eudicots	Proteaceae	<i>Peninsulapollis gillii</i> (Cookson) Dettmann y Jarzen 1988 (Fig. 4M)	
	Nothofagaceae	<i>Nothofagidites</i> sp. (Fig. 4V)	
	Ulmaceae	<i>Verrustephanoporites simplex</i> Leidelmeier 1966 (Fig. 5M–N)	
	Loranthaceae	<i>Syncolporites</i> sp. cf. <i>S. marginatus</i> Van Hoeken-Klinkenberg 1964 (Fig. 5D)	
		<i>Syncolporites</i> sp. (Fig. 5E)	
	Aquifoliaceae	<i>Gemmatricolpites subsphaericus</i> Archangelsky 1973 (Fig. 4F)	
		<i>Gemmatricolpites pergemmatus</i> Muller 1968 (Fig. 4G–H)	
	Onagraceae	<i>Diporites</i> sp. (Fig. 4J)	
	Gunneraceae	<i>Tricolpites reticulatus</i> Cookson 1947 ex Couper 1953 (Fig. 5J)	
		<i>Tricolpites</i> sp. cf. <i>T. bathyreticulatus</i> Stanley 1965 (Fig. 5H–I)	
		<i>Tricolpites</i> sp. cf. <i>T. cellatus</i> Norton in Norton and Hall 1967 (Fig. 5K)	
		<i>Tricolpites</i> sp. (Fig. 5L)	
	Verbenaceae	<i>Spinitricolpites jennerclarkei</i> Scholtz 1985 ex Jansonius and Hills 1990 (Fig. 5B)	
	Fabaceae	<i>Margocolporites</i> sp. cf. <i>M. lihokus</i> Srivastava 1972 (Fig. 4I)	
	Sapindaceae	<i>Cupanieidites</i> sp. cf. <i>C. orthoteichus</i> Cookson and Pike 1954 (Fig. 4E)	
	Haloragaceae	<i>Myriophyllumpollenites</i> sp. 1 in Quattrocchio and Volkheimer 1988 (Fig. 4K)	
	Incertae sedis		<i>Aquillapollenites magnus</i> Regali et al. 1974 (Fig. 4A–B)
			<i>Crassitricolporites brasiliensis</i> Hergreen 1972 (Fig. 4C–D)
			<i>Psilatricolpites</i> sp. (Fig. 4P)
			<i>Psilatricolporites</i> sp. 1 (Fig. 4S)
			<i>Psilatricolporites</i> sp. 2 (Fig. 4T)
			<i>Psilastephanoporites brasiliensis</i> Regali et al. 1974 (Fig. 4O)
			? <i>Psilabrevitricolpites</i> sp. (Fig. 4N)
			<i>Penetetrapites</i> sp. cf. <i>P. mollis</i> Hedlund and Norris 1968 (Fig. 4L)
			<i>Retistephanocolpites regularis</i> Hoeken-Klinkenberg 1966 (Fig. 4U)
			<i>Retitricolporites</i> sp. cf. <i>R. multibaculatus</i> Takahashi 1997 (Fig. 4W–X)
		<i>Retitricolporites</i> sp. (Fig. 4Y)	
		<i>Rousea patagonica</i> Archangelsky 1973 (Fig. 5A)	
		<i>Striatricolporites</i> sp. (Fig. 5F–G)	
		<i>Tricolporites</i> sp. 1 (Fig. 5Q)	
		<i>Tricolporites</i> sp. 2 (Fig. 5R)	
		<i>Tricolporites</i> sp. 3 (Fig. 5S)	
		<i>Tricolporites</i> sp. 4 (Fig. 5T)	
		<i>Triorites</i> sp. cf. <i>T. tenuixinus</i> Muller 1968 (Fig. 5O)	
		<i>Verrutricolpites</i> sp. (Fig. 5P)	



TRILETE SPORES

- *Apiculatisporis* sp. cf. *A. charahuillaensis* Volkheimer 1972

Diameter: 68–71 μm (6 specimens).

The original material of *A. charahuillaensis* has diameters of 35–40 μm (Volkheimer, 1972). Our specimens are larger, such as the specimens recorded in the Olmedo and Tunal formations by Narváez (2009).

Argentina. *A. charahuillaensis* was originally described for the Calovian of the Neuquén Basin (Volkheimer, 1972). The species is also recorded in the Danian of the Tunal Formation (Volkheimer et al., 2006; Narváez, 2009).

Rest of the world: not recorded.

Botanical affinity: unknown.

- *Baculatisporites comaumensis* (Cookson 1953) Potonie 1956

Diameter: 31–33 μm (2 specimens).

Worldwide distribution.

Botanical affinity. Osmundaceae.

- *Concavissimisporites* sp.

Diameter: 49–58 μm (8 specimens).

Botanical affinity. Dicksoniaceae.

- *Fisciniasporites* sp.

Diameter: 61(67)77 μm (16 specimens).

Botanical affinity. Schizaeaceae.

- *Gabonisoris vigorouxii* Boltenhagen 1967

Diameter: 41(49)58 μm (12 specimens).

Argentina: Maastrichtian–Danian and early to middle Eocene (Quattrocchio, 1978; Moroni, 1982; Limarino et al., 2000; Volkheimer et al., 2006, 2007; Narváez, 2009, Vallati et al., 2016, 2017).

Rest of the world: Bolivia: Maastrichtian (Vajda-Santivanez, 1999). Brazil: Coniacian–Maastrichtian (Regali et al., 1974b; Ashraf and Stinnesbeck, 1988). Gabón: Upper Cretaceous (Boltenhagen, 1967).

Botanical affinity. Marsileaceae (Ashraf and Stinnesbeck, 1988).

- *Gleicheniidites* sp.

Diameter: 35–42 μm (4 specimens).

Worldwide distribution.

Botanical affinity. Gleicheniaceae.

- *Microreticulatisporites* sp.

Diameter: 36(38)47 μm (10 specimens).

Botanical affinity: unknown.

- *Polypodiaceoisporites retirugatus* Muller 1968

Diameter: 41(47) 57 μm (29 specimens).

Argentina: Volkheimer et al. (2006) cites as cf. for the Tunal Formation at the type locality.
Rest of the world: *P. retirugatus* was recorded from the Late Cretaceous to Early Paleogene of Malaysia (Muller, 1968) and Neogene of Papua and New Guinea (Playford, 1982).
Botanical affinity. Polipodiaceae Family. Muller (1968) mentions the probable affinity with the extant fern genus *Pteris* L.

***Verrucosisporites* sp. 1**

Diameter: 44–66 μm (6 specimens).

Botanical affinity: unknown.

- *Verrucosisporites* sp. 2

Diameter: 60 μm (6 specimens).

Botanical affinity: unknown.

- *Zlivisporis reticulatus* (Pocock) Pacltova y Simoncsics 1970

Diameter: 48(59)68 μm (19 specimens).

Argentina: Albian (Archangelsky and Llorens, 2005)

Worldwide distribution for the Cretaceous.

Botanical affinity: Marchantiophyta.

- *Zlivisporis* sp.

Diameter: 51–62 μm (6 specimens).

Smooth proximal face, and reticulated distal face.

Botanical affinity: Marchantiophyta.

GYMNOSPERM POLLEN GRAINS

- *Podocarpidites* sp.

Diameter: 56–58 μm (2 specimens).

Worldwide distribution.

- *Equisetosporites* sp. cf. *E. concinnus* Singh 1964

Equatorial diameter: 66–88 μm , polar diameter: 20–39 μm (2 specimens).

Equisetosporites concinnus has 10–12 ribs of 7–10 μm in the thinner middle part, our specimens have less ribs (7–8).

Argentine: not previously recorded.

Rest of the world: a single record for the Maastrichtian.

Botanical affinity. Ephedraceae.

- *Equisetosporites* sp. cf. *E. notensis* (Cookson) Romero 1977

Equatorial diameter: 25–47 μm , polar diameter: 14–28 μm (6 specimens).

Argentina: Kachaiké Formation (Archangelsky and Llorens, 2009), Upper Cretaceous of the Lefipán Formation (Baldoni, 1992) and for the Eocene of the Río Turbio Formation (Romero, 1977). Middle to Upper Miocene (Barreda and Palazzesi, 2007).

Rest of the world: Cretaceous to Neogene.

Botanical affinity. Ephedraceae.

- ***Gnetaceaepollenites jansonii* (Pocock) Lima 1980**

Equatorial diameter: 75–111 μm , polar diameter: 40–42 μm (5 specimens).

Argentina: Ciénaga del Rio Huaco Formation, Maastrichtian (Limarino et al., 2000).

Rest of the world: Maastrichtian of Bolivia (Vajda-Santivanez, 1999); Albian–Cenomanian of Brasil (Herngreen, 1973; Lima, 1980); Albian of Canada (Pocock, 1964; Burden y Hills, 1989); and Aptian–Maastrichtian of Africa (Kotova, 1978; Schrank and Ibrahim, 1995).

Botanical affinity. Ephedraceae.

- ***Gnetaceaepollenites* sp. cf. *G. barghoornii* (Pocock) Lima 1980**

Equatorial diameter: 42–45 μm , polar diameter: 20–21 μm (3 specimens).

Specimens in a poor state of preservation with 8–9 ribs, *Gnetaceaepollenites barghoornii* differs in having 10–12 ribs.

Botanical affinity. Ephedraceae.

- ***Gnetaceaepollenites* sp.**

Equatorial diameter: 42 μm , polar diameter: 21 μm (1 specimens).

Botanical affinity. Ephedraceae.

- ***Singhia multicostata* (Brener) Lima 1980**

Equatorial diameter: 31–41 μm , polar diameter: 24–33 μm (9 specimens).

Lima (1980) makes a combination, considering *Ephedripites multicostatus* Brenner 1963 and *Equisetosporites multicostatus* Norris 1967 as synonyms of *Singhia multicostata*.

Argentina: Albian–Danian (Papú, 1990, 2002; Papú and Sepúlveda, 1995; Ruiz et al., 1999; Vallati, 2001; Narváez and Sabino, 2008; Narváez, 2009).

Rest of the world: Brazil: Albian (Lima, 1980). Canada: Albian–Campanian (Burden and Hills, 1989; Braman, 2001). USA: Albian (Hedlund and Norris, 1968). Gabon: Lower Cretaceous (Wood et al., 1997). Egipt: Barremian–Aptian (Sultan, 1986). Sudan: Cenomanian–Turonian (Schrank, 1994). Spain: Aptian–Albian (Arias and Doubinger, 1980).

Botanical affinity. Ephedraceae.

ANGIOSPERM POLLEN GRAINS

- ***Aquillapollenites magnus* Regali et al. 1974**

Diameter: 38(52)62 μm (16 specimens).

Argentina: Yacoraite Formation (Moroni 1982).

Rest of the world: *Tricornites elongatus* Zone, Maastrichtian of Brazil (Regali et al. 1974).

Botanical affinity. *Aquillapollenites* complex were produced by members of the Santalales Order, based on their resemblance to some extant pollen grains of the families Loranthaceae and Santalaceae (Jarzen, 1977, Muller, 1984).

- ***Arecipites* sp. cf. *A. tenuixinus* Leffingwell 1971**

Diameter: 29–32 μm (8 specimens).

According to the original description it has 0.5 μm lumina and a dark exine, in our specimens the lumina are smaller than 1 μm .

Argentina: not previously recorded.

Rest of the world: Paleocene (Leffingwell, 1971), Maastrichtian (Nichols, 2002), Paleocene (Nichols and Brown, 1992), Santonian–Campanian (Braman, 2001).

Botanical affinity: Arecaceae.

- *Crassitricolporites brasiliensis* Herngreen 1972

Diameter: 35–53 μm (7 specimens).

Argentina: Yacoraite Formation (Moroni, 1982).

Rest of the world: Senonian of Brazil (Herngreen, 1972), *Crassitricolporites brasiliensis* Zone, Maastrichtian (Regali et al., 1974).

Botanical affinity: unknown.

- *Cupanieidites* sp. cf. *C. orthoteichus* Cookson and Pike 1954

Diameter: 25 μm (3 specimens).

Argentina: not previously recorded.

Rest of the world: Maastrichtian of Antarctica (Askin, 1990)

Botanical affinity: Sapindaceae, Cupaniae tribe.

- *Diporites* sp.

Diameter: 44 μm (1 specimens).

Botanical affinity: Onagraceae Family.

- *Gemmatricolpites subsphaericus* Archangelsky 1973

Diameter: 22–32 μm (2 specimens).

Argentina: Salta Basin: Mealla Formation (Selandian, Quattrocchio et al., 1997); Maíz Gordo Formation (Tanetiano, Quattrocchio and del Papa, 2000); Tunal Formation (Daniano, Quattrocchio and Volkheimer, 1988; Volkheimer et al., 2006). Patagonia: Salamanca Formation (Paleoceno, Archangelsky, 1973); Cerro Bororó Formation (Danian, Volkheimer et al., 2007).

Rest of the world: not previously recorded.

Botanical affinity: Aquifoliaceae.

- *Gemmatricolpites pergemmatum* Muller 1968

Diameter: 20–21 μm (2 specimens).

Argentina: not previously recorded.

Rest of the world: The species has its first record from the Senonian to Eocene sediments of Sarawak (Muller, 1968) and later from the Campanian–Maastrichtian of Egypt (Schrank, 1987) and from late Maastrichtian–Paleocene of Sudan (Awad, 1994; Eisawi and Schrank, 2008; Eisawi, 2015).

Botanical affinity: Aquifoliaceae, *Ilex* sp. (Muller, 1968).

- *Liliacidites variegatus* Couper 1953

Diameter: 28(30)44 μm (12 specimens).

Argentina and rest of the world: widely recorded since Cenomanian.

Botanical affinity: Liliaceae Family.

- ***Margocolporites* sp. cf. *M. lihokus* Srivastava 1972**

Diameter: 30 μm (3 specimens), colpi long, with 1.5 μm thick margo. Exine reticulate, meshes subcircular, ranging from 0.3–0.7 μm .

Margocolporites lihokus have larger diameter and thicker margo.

Argentina: not previously recorded.

Rest of the world: Paleocene, USA (Elsik, 1968; Srivastava, 1972), Maastrichtian (Farabee and Canright, 1986; Nichols 2002).

Botanical affinity: Fabaceae.

- ***Myriophyllumpollenites* sp. in Quattrocchio and Volkheimer 1988**

Diameter: 24(27)29 μm (10 specimens).

Argentina: Salta Basin: Tunal and Olmedo formations (Danian; Quattrocchio and Volkheimer, 1988; Volkheimer et al., 2006; Narvaez, 2009), Mealla Formation (middle Paleocene, Quattrocchio et al., 1997).

Rest of the world: not previously recorded.

Botanical affinity: Haloragaceae.

- ***Nothofagidites* sp.**

Diameter: 26–30 μm (4 specimens).

Botanical affinity: Nothofagaceae.

- ***Penetetrapites* sp. cf. *P. mollis* Hedlund and Norris 1968**

Diameter: 15–24 μm (2 specimens).

Argentina: Santonian of the Río Guanaco Formation (Pothe de Baldi, 1995).

Rest of the world: Albian–Campanian.

Botanical affinity: unknown.

- ***Peninsulapollis gillii* (Cookson) Dettmann and Jarzen 1988**

Diameter: 25 μm (1 specimens).

Argentina: Late Cretaceous of Chubut Province (Papu, 1989; Baldoni and Askin, 1993); Maastrichtian from Mendoza Province (Papú, 2002); Late Cretaceous of Mendoza Province (Papú, 1990, 2002); Maastrichtian–Paleocene of the Colorado Basin (Ruiz and Quattrocchio, 1997; Quattrocchio and Ruiz, 1999); Maastrichtian of Santa Cruz Province (Povilauskas et al., 2008); Paleocene–Danian of Chubut Province (Archangelsky, 1973; Volkheimer et al., 2007); Paleocene–Eocene (Baldoni, 1987); Eocene of Santa Cruz Province (Romero and Zamaloa, 1985); Eocene of Santa Cruz Province (Romero and Castro, 1986), Campanian of Río Negro Province (Papú and Sepúlveda, 1995), Paleogene of Río Negro and Neuquén provinces (Melendi et al., 2003), Campanian–Maastrichtian of Salta Province (Narvéez and Sabino, 2008).

Rest of the world: Early Eocene and Late Eocene–Early Oligocene (Fasola, 1969). Antarctica: Late Campanian–Paleocene (Baldoni and Barreda, 1986; Cao, 1992; Dolding, 1992; Amenábar and Di Pasquo, 2008); Santonian–Paleogene (Askin and Baldoni, 1998); Eocene and Oligocene (Mohr, 1990); Cenozoic (Bratzeva, 1983). Australia: Campanian and Early Eocene (Dettmann and Jarzen, 1988); Late Cretaceous to Early Eocene (Cookson, 1957). New Zealand: Late Cretaceous to Neogene (Mc Intyre, 1968, Raine et al., 2008).

Botanical affinity: Proteaceae.

- **?*Psilabrevitricolpites* sp.**

Diameter: 38 μm (3 specimens), short (8–10 μm) and thickened (2,8 μm) colpi.
Botanical affinity: unknown.

- ***Psilastephanoporites brasiliensis* Regali et al. 1974**

Diameter: 29–32 μm (7 specimens).
Argentina: as cf., in the Yacoraite Formation (Moroni 1982)
Rest of the world: Cenomanian of Brazil (Jaramillo et al., 2010); Santonian of Brazil (Lima, 1981; Muller, 1987).
Botanical affinity: unknown.

- ***Psilatricolpites* sp.**

Diameter: 35–42 μm (3 specimens).
Botanical affinity: unknown.

- ***Psilatricolporites* sp. 1**

Diameter: 37 μm (1 specimen), exine 1–1.2 μm . Subcircular amb.
Botanical affinity: unknown.

- ***Psilatricolporites* sp. 2**

Diameter: 19 μm (1 specimen). Subtriangular amb.
Botanical affinity: unknown.

- ***Retistephanocolpites regularis* Hoeken-Klinkenberg 1966**

Diameter: 29–37 μm (8 specimens).
Argentina: Campanian to Early Maastrichtian, Salta Basin (Quattrocchio et al., 2005)
Rest of the world: Middle Eocene of Nigeria (Van Hoeken-Klinkenberg, 1964), Paleogene (Zaklinskaja, 1978), Paleocene–Early Eocene (Bio-lokoto et al., 1998).
Botanical affinity: unknown.

- ***Retitricolporites* sp. cf. *R. multibaculates* Takahashi 1997**

Diameter: 46–57 μm (3 specimens); muri 0,9–1,1 μm , lumina 1.7–5 μm irregular to subpolygonal with a loose columella on the floor of many lumina; exine 2.3–3 μm .
Argentine: not previously recorded.
Rest of the world: Campanian of Russia (Takahashi, 1997)
Botanical affinity: unknown.

- ***Retitricolporites* sp.**

Diameter: 28 μm ; muri 0,5–0,6 μm , lumina 0.5–1.2 μm , muri wider than lumina (except at elongated lumina) (1 specimens).
Botanical affinity: unknown.

- ***Rousea patagonica* Archangelsky 1973**

Diameter: 31–45 μm (8 specimens).

Argentina: Maastrichtian to Paleocene (Archangelsky, 1973; Papu, 1989 (as cf.), Baldoni and Askin, 1993; Quattrocchio et al., 1997, 1999, 2000; Prámparo and Papú, 2003; Volkheimer et al., 2006, 2007; Narváz and Sabino, 2008; Narváz. 2009, Povilaukas, 2013).

Rest of the world: Eocene (Kedves, 1986).

Botanical affinity: unknown.

- ***Spinitricolpites jennerclarkei* Scholtz 1985 ex Jansonius y Hills 1990**

Diameter: 50–53 μm (2 specimens).

Argentina: Los Blanquitos and Lecho formations (Late Senonian) (Quattrocchio et al., 2005).

Rest of the world: Paleocene of South Africa (Scholtz 1985).

Botanical affinity: Verbenaceae (Quattrocchio et al., 2005).

- ***Spinizonocolpites* sp.**

Diameter: 42 μm , broad-based spines (1.7–2.1 μm) and 3–3.5 μm long (1 specimens).

Botanical affinity: Arecaceae.

- ***Syncolporites* sp. cf. *S. marginatus* Van Hoeken-Klinkenberg 1964**

Diameter: 18–24 μm (2 specimens).

Argentine: not previously recorded.

Rest of the world: Maastrichtian of Sudan (Awad, 1994), Egypt (Schrank, 1984), and Nigeria (Van Hoeken-Klinkenberg 1964), Assemblage Zone K-IV Campanian–Maastrichtian (Cole et al., 2017), Assemblage Zone V Maastrichtian of Sudan (Eisawi et al., 2012), Paleogene of Colombia (Jaramillo and Dilcher, 2001).

Botanical affinity: Loranthaceae.

- ***Syncolporites* sp.**

Diameter: 18 μm (1 specimens).

Botanical affinity: Loranthaceae.

- ***Striatricolporites* sp.**

Diameter: 38–42 μm (16 specimens), tricolporate, subprolate to prolate, regulate Exine, colpi long thickened, ora discernible.

Comparison: *Striatricolporites gamerroi* Archangelsky 1973 presents smaller diameter and exine microreticulate with stries development in polar way.

Botanical affinity: unknown.

- ***Tricolpites reticulatus* Cookson 1947**

Diameter: 24(32)44 μm (23 specimens).

Argentina: Late Campanian–Miocene (Archangelsky, 1973; Quattrocchio, 1978a, 1980b; Romero and Zamaloa, 1985; Guerstein, 1990; Baldoni and Askin, 1993; Papú and Sepúlveda, 1995; Prámparo et al., 1995, 1996; Barreda, 1996, 1997; Quattrocchio et al., 1997; Ruiz and Quattrocchio, 1997; Quattrocchio and Ruiz, 1999; Ruiz et al., 1999; Barreda and Palamarczuk, 2000; Palamarczuk and Barreda, 2000; Papú, 2002; Melendi et al., 2003; Volkheimer et al., 2006, 2007; Barreda and Palazzesi, 2007; Prámparo et al., 2007).

Rest of the world: Antarctica: Middle Miocene (Cookson, 1947; Cranwell, 1969). Chile: Eocene–Early Oligocene (Fasola, 1969). Canada: Late Cretaceous (Jarzen and Norris, 1975). New Zealand: Late

Oligocene–Early Miocene (Pocknall and Mildenhall, 1984). Indian Ocean: Middle Paleocene (Harris, 1974); Late Oligocene (Kemp, 1974); Middle Cenozoic (Kemp and Harris, 1977). New Guinea: Pliocene (Cookson and Pike, 1954).

Botanical affinity: Gunneraceae (*Gunnera* sp.).

- ***Tricolpites* sp. cf. *T. bathyreticulatus* Stanley 1965**

Diameter: 30–46 μm (4 specimens).

It differs by having a larger diameter than *T. bathyreticulatus*.

Argentina: Eocene Río Turbio Formation, Santa Cruz Province (Fernández, 2018).

Rest of the world: widespread species widely distributed in Cenomanian to early Miocene deposits (see citations in White, 2006 and Raine et al., 2011).

Botanical affinity: Gunneraceae.

- ***Tricolpites* sp. cf. *T. cellatus* Norton in Norton and Hall 1967**

Diameter: 44–55 μm (6 specimens), muri 0.4–0.7 μm , irregular lumina up to 1 μm and elongated ones up to 1.5 μm , exine 0.9–1 μm .

Tricolpites cellatus has a smaller diameter (32–40 μm), and an exine of 1.5 μm .

Argentina: not previously recorded.

Rest of the world: Upper Cretaceous–Lower Paleogene of USA (Norton and Hall 1967, 1969), K–Pg of Canada (Jerzykiewick and Sweet, 1986).

Botanical affinity: Gunneraceae.

- ***Tricolpites* sp. 1**

Diameter: 31–45 μm (2 specimens), muri 0.8–1 μm , lumina 0.8–3 μm irregular, often elongated, exine 1.5–1.8 μm .

Retitricolpites virgeus presents smaller diameter and size of luminae.

Botanical affinity: Gunneraceae.

- ***Tricolporites* sp. 1**

Diameter: 25–31 μm (3 specimens), muri 0.2–0.3 μm , lumina 0.2–0.4 μm , colpi and pore thickened.

Botanical affinity: unknown.

- ***Tricolporites* sp. 2**

Diameter: 23–27 μm (3 specimens), muri 0.3 μm , subpolygonal to irregular lumina 0.2–1.4 μm (3 specimens).

Botanical affinity: unknown.

- ***Tricolporites* sp. 3**

Diameter: 29–39 μm (2 specimens), muri 0.2–0.3 μm , subpolygonal to irregular lumina 0.3–0.5 μm (2 specimens).

Botanical affinity: unknown.

- ***Tricolporites* sp. 4**

Diameter: 35–46 μm (4 specimens), muri 0.3–0.4 μm , lumina 0.3–1 μm , exine 0.8–0.9 μm , colpi membrane with grains of 0.8–0.9 μm .

Botanical affinity: unknown.

- ***Triorites* sp. cf. *T. tenuixinus* Muller 1968**

Diameter: 35 μm (1 specimens).

A single specimen not well preserved.

Argentina: not previously recorded.

Rest of the world: Cenomanian of Nigeria (Oloto and Yikarebogha, 2013), (Muller 1968), Paleocene (Salard-Cheboldaeff, 1979).

Botanical affinity: unknown.

- ***Verrustephanoporites simplex* Leidelmeyer 1966**

Diameter: 28(25)35 μm (14 specimens).

Argentina: Salta Basin: transitional strata of Los Blanquitos and Lecho formations (Late Senonian)

(Quattrocchio et al., 2005), Tunal Formation (Danian, Quattrocchio and Volkheimer, 1988;

Volkheimer et al., 2006), Olmedo Formation (Danian, Narváez, 2009), Mealla Formation (Selandian,

Quattrocchio et al., 1997) and Maíz Gordo Formation (Tanetian) (Volkheimer et al., 1984); Colorado

Basin: Pedro Luro Formation (Paleocene) (Ruiz and Quattrocchio, 1997; Quattrocchio and Ruiz,

1999); Neuquén Basin: Jagüel Formation (upper Maastrichtian; Prámparo y Papú, 2003). Anta

Formation (Early/Mid Miocene, Quattrocchio et al., 2003).

Rest of the world: Paleocene of Guyana (Leidelmeyer, 1966).

Botanical affinity: Ulmaceae (*Phyllostylon* sp.).

- ***Verrutricolpites* sp.**

Diameter: 46 μm , exine 2 μm , warts 1.4–2 μm wide at base and 0.9–1.6 μm high (1 specimens).

Botanical affinity: unknown.

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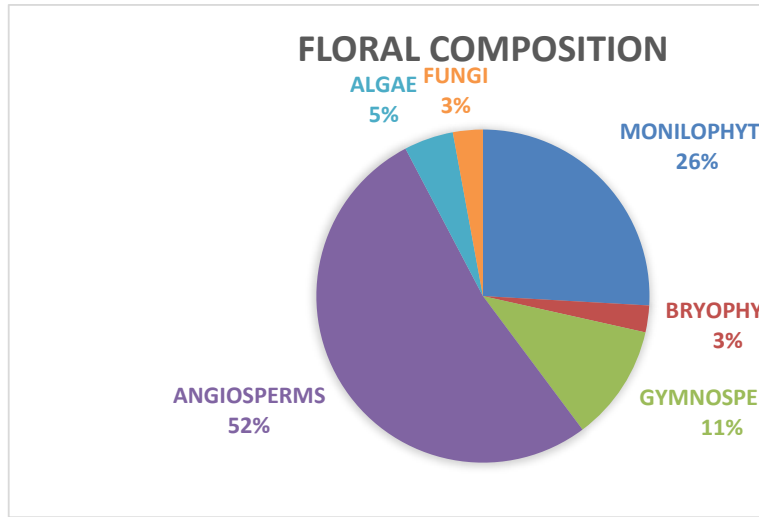
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MONIOPHYTA	25.90%
BRYOPHYTA	2.60%
GYMNOSPERMS	11.30%
ANGIOSPERMS	52.50%
ALGAE	4.80%
FUNGI	2.90%
	100.00%





A

TA

RMS



Tci (Briophyta)	2.60%
TA	3.7
TT	2.5
TS	3.9
TP (<i>Gabonisorites</i> sp.)	1.8
TC (<i>Polypodiaceoisorites</i> sp.)	14
Total Monylophyta	25.90%
Fungi	2.90%
Ovoidites sp.	4.80%



PP
2S

10.30%
1.00%





3C	20.50%
3P	2.60%
3CP	7%
1C	9.10%
4P	13.20%
	52.40%





	Pollen grains	Taxon	
3C	Tricolpated		Tci
3P	Triplicated		TA
3CP	Tricolporated		TT
1C	Monocolpated		TL
4P	Tetraplicated	<i>Verrustephanocolpites</i> sp.	TP
PP	Poliplicated		TC
2S	Bisacated		





Spores	Taxon
Trilete cingulated	Briophyta
Trilete apiculated	
Trilete tricassated	
Trilete smooth	
Trilete perisporated	<i>Gabonisorites</i> sp.
Trilete cingulated	<i>Polypodiaceoisporites</i> sp.





Conflict of Interest: none

