

## RESEARCH/REVIEW ARTICLE

# Variability of nutrient and thermal structure in surface waters between New Zealand and Antarctica, October 2004–January 2005

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## Abstract

We describe the upper ocean thermal structure and surface nutrient concentrations between New Zealand and Antarctica along five transects that cross the Subantarctic Front, the Polar Front (PF) and the southern Antarctic Circumpolar Current (ACC) front. The surface water thermal structure is coupled with variations in surface nutrient concentrations, making water masses identifiable by both temperature and nutrient ranges. In particular, a strong latitudinal gradient in orthosilicate concentration is centred at the PF. On the earlier sections that extend south-west from the Campbell Plateau, orthosilicate increases sharply southward from 10–15 to 50–55  $\mu\text{mol l}^{-1}$  between 58° S and 60° S, while surface temperature drops from 7°C to 2°C. Nitrate increases more regularly toward the south, with concentrations ranging from 10–12  $\mu\text{mol l}^{-1}$  at 54° S to 25–30  $\mu\text{mol l}^{-1}$  at 66° S. The same features are observed during the later transects between New Zealand and the Ross Sea, but the sharp silica and surface temperature gradients are shifted between 60° S and 64° S. Both temporal and spatial factors may influence the observed variability. The January transect suggests an uptake of silica, orthophosphate and nitrate between 63° S and 70° S over the intervening month, with an average depletion near 37%, 44% and 29%, respectively. An N/P (nitrite+nitrate/orthophosphate) apparent drawdown ratio of  $8.8 \pm 4.1$  and an Si/N (silicic acid/nitrite+nitrate) apparent drawdown ratio  $> 1$  suggest this depletion results from a seasonal diatom bloom. A southward movement of the oceanic fronts between New Zealand and the Ross Sea relative to prior measurements is consistent with reports of recent warming and changes in the ACC.

The Antarctic Circumpolar Current (ACC) is the largest coherent current structure in the world's oceans. A pronounced meridional gradient in surface properties that separate waters of the Southern Ocean from the warmer and saltier waters to the north is called the Subtropical Front (STF) (Clifford 1983; Hofmann 1985). South of the STF is the strong eastward flow of the ACC, the primary means by which water is exchanged between the major ocean basins. In particular, the ACC plays an essential role in the thermohaline circulation and the

global redistribution of heat and other properties (Rintoul & Bullister 1999). Deacon (1937) was the first to note that the southward transition from warm, light subtropical water occurred in a step-like manner, rather than as a gradual change across the breadth of the Southern Ocean. These steps are associated with strong surface currents (Nowlin et al. 1997), characterized by two or three narrow jets associated with sharp fronts (Nowlin & Klinck 1986; Orsi et al. 1995; Belkin & Gordon 1996; Sokolov & Rintoul 2009a). The fronts extend through the

upper ocean and usually separate zones of relatively uniform water properties (Budillon & Rintoul 2003). This paper describes the evolution of nutrient concentrations and thermal structures in the surface waters between New Zealand and Antarctica, using data collected during five transects completed from mid-October 2004 into early January 2005.

In the region south of New Zealand the zonal orientation and latitudinal separation of the physical and chemical fronts minimize the effects of convergence and divergence. The sections cross the Subantarctic Front (SAF), the Polar Front (PF) and the southern Antarctic Circumpolar Current front (SACCF) (Whitworth 1980; Russo et al. 1999; Chaigneau & Morrow 2002; Budillon & Rintoul 2003). A variety of criteria has been used to identify the fronts (Peterson & Stramma 1991; Belkin & Gordon 1996). Here we use definitions based on temperature since we rely on expendable bathythermograph (XBT) sections to identify the fronts. The SAF is identified by the maximum gradient in the range of 3–8°C at the depth of 300 m (Belkin & Gordon 1996). Its northern part (NSAF) is associated with a maximum temperature gradient in the range of 4–7°C at 300 m and its southern part (SSAF) with a maximum temperature gradient in the range of 3–4°C at 300 m (Rintoul et al. 1997). The PF is identified where temperatures to the south are colder than 2°C at 200 m (Orsi et al. 1995). The SACCF displays temperatures to the south below 0°C at depths lower than 150 m and to the north above 1.8°C at depths > 500 m (Orsi et al. 1995).

The sea floor underneath the ACC is the most important area where biogenic silica accumulates in the world ocean (DeMaster 1981). Although relatively rich in major nutrients, most of the ACC is less than optimally productive (Jaques 1989; Banse 1996), rendering it one of the most prominent high-nutrient, low-chlorophyll regions. In the South Atlantic sector it has been estimated that the total productivity of organic matter is low, which would imply a rather low drawdown of carbon dioxide (CO<sub>2</sub>) by phytoplankton (Arrigo et al. 1999; Hoppema et al. 2000). However, the Polar Front Zone (PFZ), located between the SAF and the PF (Orsi et al. 1995), is a unique oceanographic environment exhibiting some of the highest phytoplankton biomass and primary productivity in the Southern Ocean (De Baar et al. 1995). The PF separates the Antarctic Zone to the south, where high winter silicic acid concentrations result from upwelling of Circumpolar Deep Water (CDW) and from the PFZ to the north, where silica-depleted Antarctic Intermediate Water sinks and begins its northward transit (Morrison et al. 2001).

Most similar ACC studies have focused on the physical and chemical conditions encountered during a single cruise or temporal changes between data collected in different years and seasons. Morrison et al. (2001) show that the evolution of water properties is consistent with the seasonal biological cycle, so that data collected at one location represent temporal variability within a dynamic zonal current. Here we describe nitrate, silicic acid and orthophosphate levels in the surface layer and temperatures in the upper 700 m of the ACC south of New Zealand. The aim of this work is to determine the spatial and temporal variability of the physical and chemical properties of the major ocean fronts in that region during spring and summer.

## Methods

Hydrographic data were obtained during cruises aboard the RV *N.B. Palmer* (during the Antarctic Slope project) and the RV *Italica* (Table 1). Two of the transects are located between the Campbell Plateau and George V Land and three between New Zealand and the Western Ross Sea (Fig. 1). Vertical temperature profiles were obtained with T4 and T7 XBTs (Lockheed Martin Sippican, Marion, MA, USA) deployed to maximum depths of 460 m and 760 m, respectively. Seawater samples were collected at the surface underway from the ship's clean seawater flow-through systems, keyed to XBT launches spaced ca. 15' in latitude (30' in latitude during the third transect; Table 1). The XBT profiles recorded at a resolution of 0.7 m were despiked and averaged at 5-m depth intervals. The XBT accuracy is ca. 0.1°C. Sea surface temperature and salinity were measured continuously by a thermosalinograph connected to the ship's seawater intake line.

Most nutrient samples from RV *N.B. Palmer* were filtered (GF/F grade 25-mm glass microfibre filter with a nominal pore size of 0.7 µm; Whatman, Maidstone, Kent, UK), stored at –70°C in polyethylene vials and analysed on board as soon as possible after collection. However, orthophosphate samples were analysed in our Italian laboratory six months after collection because of technical problems on board. Nutrient samples from RV *Italica* were similarly filtered and analysed ashore four months later, after storage at –20°C in polyethylene vials. The nutrients (nitrite [NO<sub>2</sub>], nitrate [NO<sub>3</sub>], orthophosphate [PO<sub>4</sub>] and silicic acid-Si[OH]<sub>4</sub>) were analysed with an AutoAnalyzer trAAcs 800 system (Technicon Corporation, New York, NY, USA) and the resulting data processed with AACE<sup>®</sup> 5.24 software. Nutrient concentrations were determined employing a modification of procedures developed by Strickland & Parsons (1972).

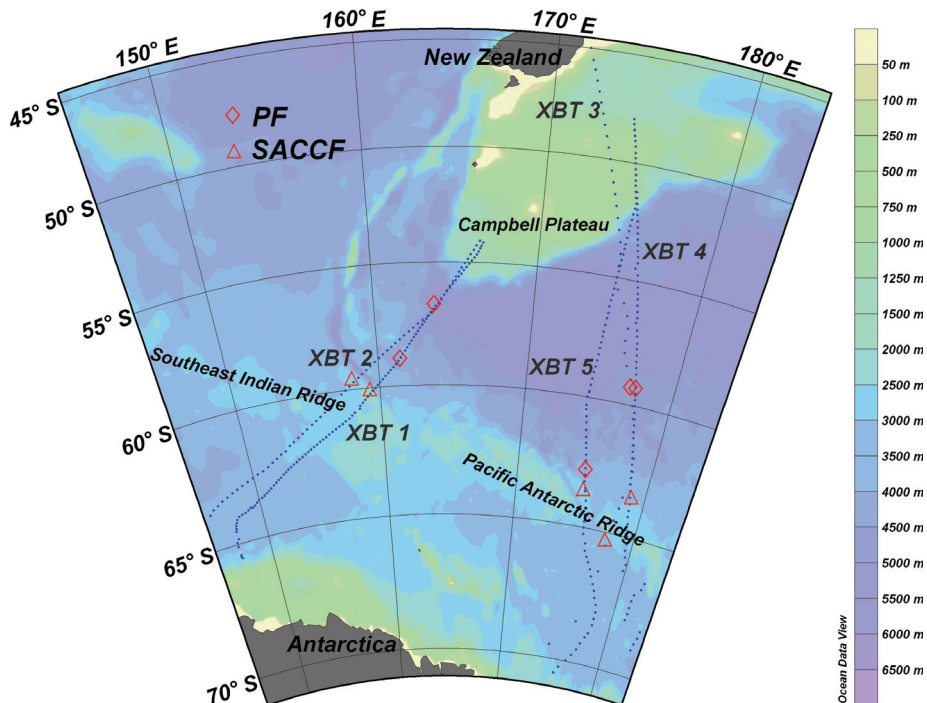
**Table 1** Summary of the expendable bathythermograph (XBT) stations and seawater samples collected at each transect during the Antarctic Slope (AnSlope3) cruise and the RV *Italica* (ITA05) cruise in 2005.

Transects	Period	Number of XBT stations	Number of seawater samples
AnSlope3-XBT1	15–20 October 2004	108	108
AnSlope3-XBT2	1–4 November 2004	67	67
AnSlope3-XBT3	7–15 November 2004	78	45
AnSlope3-XBT4	30 November–5 December 2004	74	72
ITA05-XBT5	1–6 January 2005	96	89

For nitrite ( $\text{NO}_2$ ) and nitrate ( $\text{NO}_3$ ) we used the procedure whereby nitrate is reduced to nitrite at pH 8 in a copper–cadmium reductor. The nitrite then reacts under acid conditions with sulfanilamide to form a diazo compound that then couples with *N*-(1-naphthyl)-ethylenediamine, to form a reddish-purple azo dye that is measured at 550 nm. Nitrate is corrected for nitrite contribution by correcting for the efficiency of nitrate reduction and subtracting the nitrite concentration measured in an unreduced portion of the sample. Here we consider only nitrite plus nitrate ( $\text{NO}_2 + \text{NO}_3$ ) and the coefficient of variability is 0.44%. The determination of orthophosphate ( $\text{PO}_4$ ) is based on the colourimetric method in which a blue compound is formed by the reaction of phosphate, molybdate and antimony followed by reduction with ascorbic acid. The reduced blue phosphor–molybdenum complex is read at 880 nm,

with a coefficient of variability of 0.57%. Measurement of soluble silicic acid ( $\text{Si}[\text{OH}]_4$ ) is based on the reduction of a silico molybdate in acid solution to molybdenum blue by ascorbic acid. Oxalic acid is introduced to the sample to minimize interference from orthophosphate and the absorbance is measured at 660 nm with a coefficient of variability of 0.59%.

Estimates of surface chl-*a* concentration (level 3, eight-day composites) were derived from data from the National Aeronautics and Space Administration (NASA) Sea-viewing Wide Field-of-view Sensor (SeaWiFS) project. Data were obtained from the Goddard Distributed Active Archive Center using the standard OC4v5.2 chlorophyll algorithm (<http://oceancolor.gsfc.nasa.gov/REPROCESSING/SeaWiFS/R5.2/>). Estimates of sea surface temperature derived from level 3, eight-day composite images from the Moderate Resolution Imaging

**Fig. 1** Geographic locations of five expendable bathythermograph (XBT) sections between New Zealand and Antarctica made by the research vessels *N.B. Palmer* and *Italica* from mid-October 2004 into early January 2005. The positions of main fronts—the Polar Front (PF) and the southern Antarctic Circumpolar Current front (SACCF)—are plotted.

Spectroradiometer launched on the Aqua satellite have been obtained from the Goddard Distributed Active Archive Center using a K-490v1.1 algorithm (<http://oceancolor.gsfc.nasa.gov/REPROCESSING/Aqua/R1.1/>). Finally, diatom surface concentrations were obtained from NASA's Ocean Biogeochemical Model (NOBM; <http://reason.gsfc.nasa.gov/Giovanni/>) (Gregg 2008).

## Results

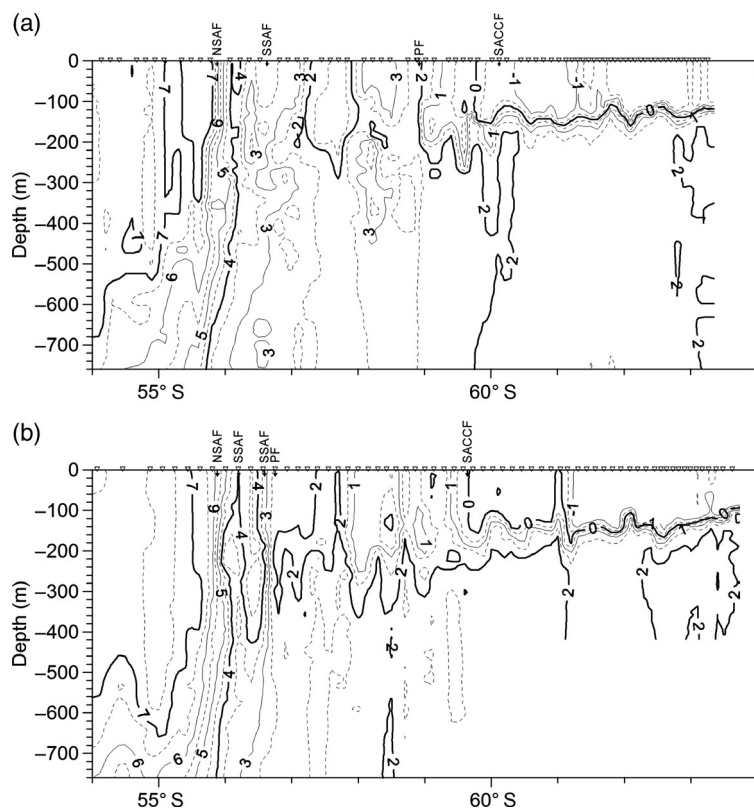
### Thermal structure and variability

Figure 2 shows the upper ocean temperature field in mid-October and early November 2004 between the southwest end of the Campbell Plateau and George V Land. In October the SAF is split into a NSAF (55.9° S) and a SSAF (56.6° S). The PF is centred at 58.9° S and the SACCF at 60.1° S. A cold core with PF characteristics evident to the north at 57.5° S could be an eddy or meander.

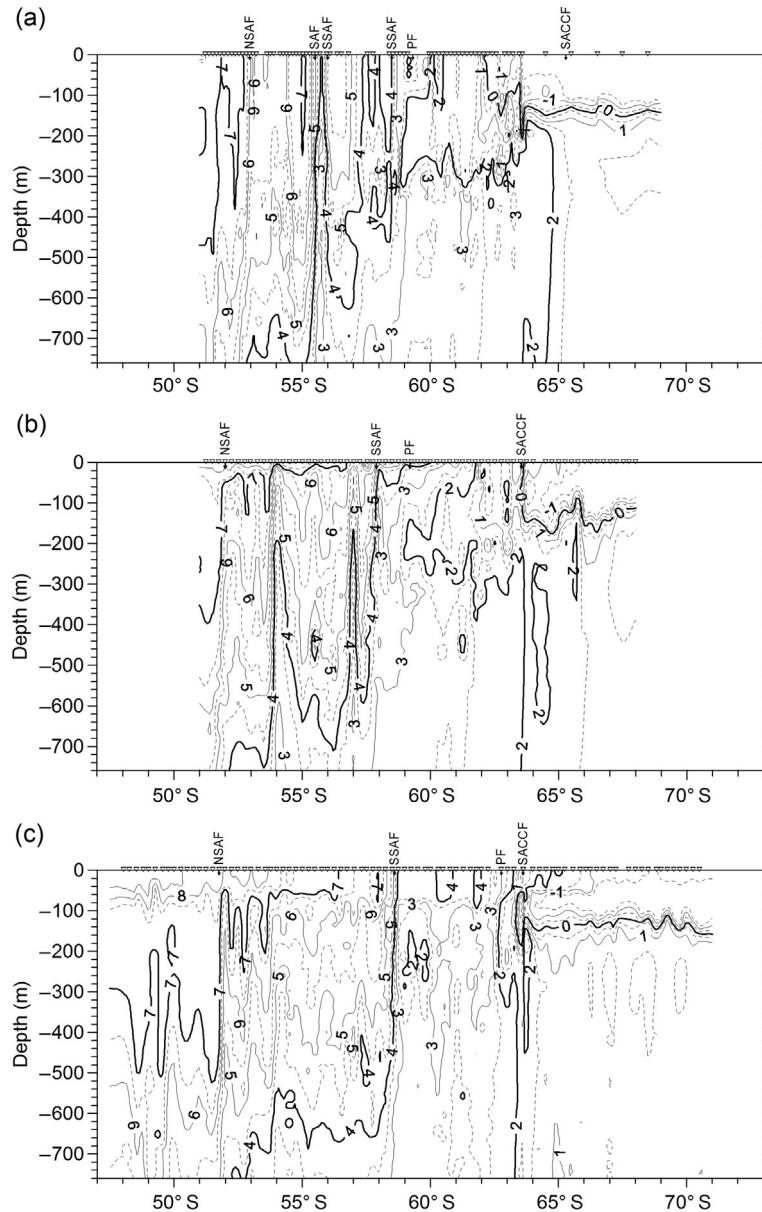
Less than three weeks later, the November section is characterized again by an intense NSAF centred at 55.9° S, with temperatures decreasing from 7°C to 5°C in

< 80 km. The SSAFs can also be seen both at 56.2° S and at 56.6° S, suggesting the presence of a cold eddy/meander in this area, as observed by Budillon & Rintoul (2003). The PF has shifted to the north (56.7° S), nearly coinciding with the southern SSAF. The SACCF is also shifted to the north, ca. 0.5° of latitude, compared to the previous period.

Figure 3 shows a sequence of three sections taken in mid-November and early December 2004 and in early January 2005 between New Zealand and the Ross Sea. In November, before summer surface heating had begun, the SAF appears at 55.5° S but also displays a northern branch at 52.9° S. Southern branches of the SAF are detected at 56° S as 58.5° S, suggesting the presence of a cold meander. The SAF, NSAF and SSAF extend to the sea surface where fronts also appear in sea surface temperature. The PF is located at 59.3° S and the SACCF at 65.3° S. By the beginning of December (Fig. 3b) the NSAF has shifted northward and is located at 52° S. The SSAF has shifted from 58.5° S to 57.9° S and the cold eddy/meander is not present. No fronts are present at the surface, where their expression appears to be eroded by the initial summer warming. The PF is in about the same



**Fig. 2** Temperature sections between New Zealand and George V Land during (a) 15–20 October 2004 and (b) 1–4 November 2004. Expendable bathythermograph locations are indicated above the plot. Positions of the northern Subantarctic Front (NSAF), the southern Subantarctic Front (SSAF), the Polar Front (PF) and the southern Antarctic Circumpolar Current front (SACCF) are also shown. Latitudinal ranges: 54° S–64° S.



**Fig. 3** Temperature sections between New Zealand and the western Ross Sea during (a) 7–15 November 2004, (b) 30 November–5 December 2004 and (c) 1–6 January 2005. Station locations are indicated above the plot. Positions of the Subantarctic Front (SAF), the northern Subantarctic Front (NSAF), the southern Subantarctic Front (SSAF), the Polar Front (PF) and the southern Antarctic Circumpolar Current front (SACCf) are also shown. Latitudinal ranges: 47° S–73° S.

position as in November, while the SACCf has moved north to 63.5° S, by 1.8°.

In early January (Fig. 3c) the NSAF (51.7° S) has moved slightly northward, the SSAF southward to 58.6° S and the PF has shifted southward to 62.8° S closer to the SACCf (63.6° S). These positions are within the ranges reported by Budillon & Rintoul (2003) for 1994–2001. The PF tracks sea floor topography into the Pacific

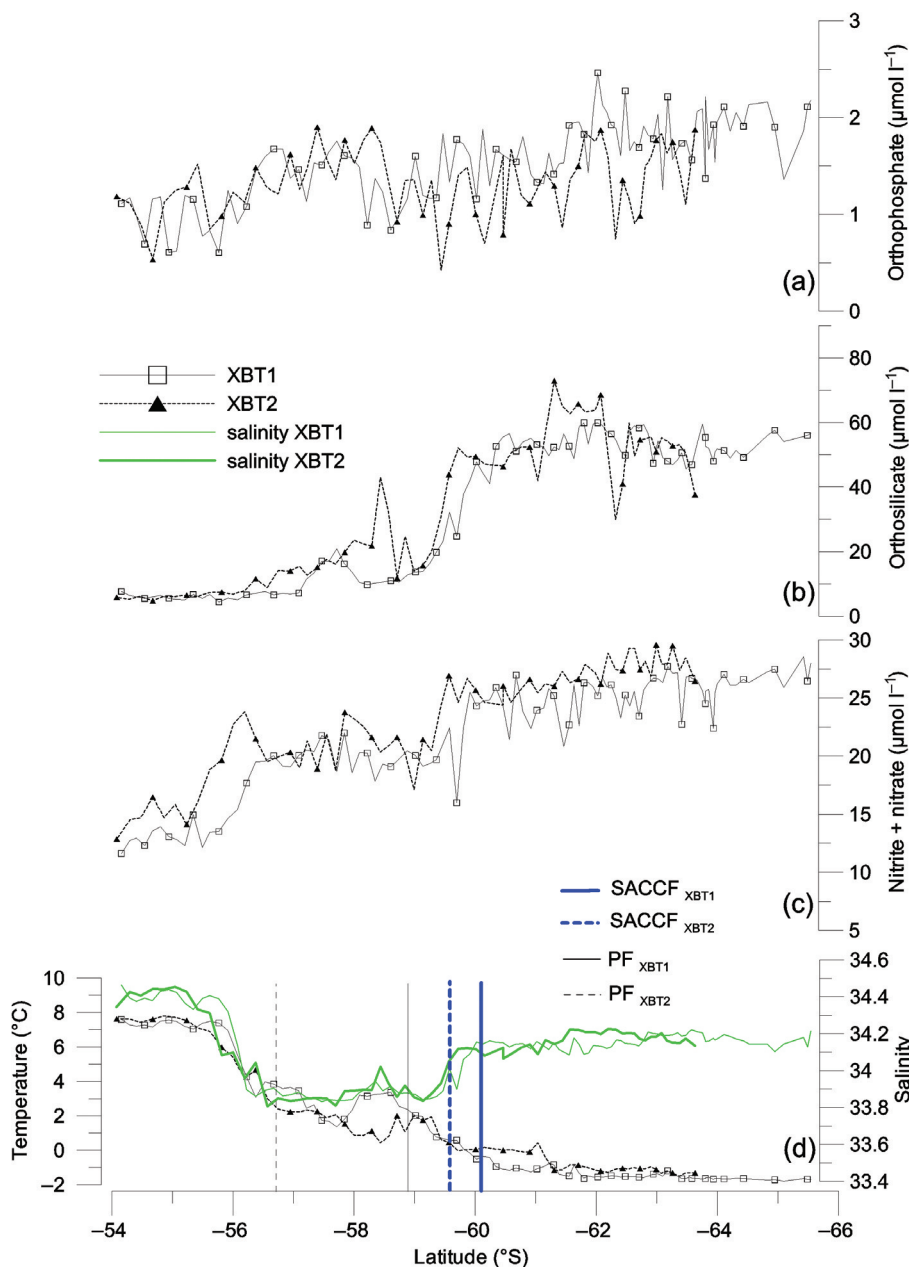
and is intensified at major bathymetric features, including the Pacific–Antarctic Ridge and Southeast Indian Ridge (Moore et al. 1999). They found a lower seasonal variability where topographic steering of the PF is strong and more seasonality from 175° E to 170° W, where the PF moved southward during the summer and farthest north during winter. While their seasonal variations in PF properties were relatively small compared with the

spatial variability, the PF mean width and temperature gradient were highest during the spring, as in this study.

### Surface variations in nutrients

During the period of study, surface water temperature and salinity change is coupled with variations in nutrient

concentrations. Figure 4 shows the surface concentrations of nutrients, temperature and salinity on the XBT 1 and XBT 2 sections between New Zealand and George V Land. In October orthophosphates increase irregularly southward, ranging from 0.8–1.1  $\mu\text{mol l}^{-1}$  at 54° S to 1.9–2.2  $\mu\text{mol l}^{-1}$  at 65.5° S. Nitrate concentrations range between 12–13  $\mu\text{mol l}^{-1}$  at 54° S and 24–28  $\mu\text{mol l}^{-1}$  at



**Fig. 4** Distribution of surface temperature (°C), salinity (green line) and nutrients (silicic acid, orthophosphate and nitrite+nitrate [µM]) between New Zealand and George V Land from 15 to 20 October 2004 (XBT 1) and from 1 to 4 November 2004 (XBT 2). (a) Orthophosphate, (b) orthosilicate and (c) nitrite+nitrate. The symbols represent every third sampling event. In (d) the vertical lines represent the Polar Front (PF) and the southern Antarctic Circumpolar Current front (SACCF) positions along the transects.

65.5° S, with two abrupt increases: the first from 12.2  $\mu\text{mol l}^{-1}$  to 19.5  $\mu\text{mol l}^{-1}$  is observed between 55.5° S and 56.4° S as temperature decreases from 7.4°C to 3.1°C across the NSAF. A further sharp increase in nitrate concentrations occurs as temperature decreases between 58.6° S ( $T=3.4^\circ\text{C}$ ;  $\text{NO}_3=19.1 \mu\text{mol l}^{-1}$ ) and 61.8° S ( $T=-1.6^\circ\text{C}$ ;  $\text{NO}_3=26.2 \mu\text{mol l}^{-1}$ ) where the PF and the SACCF are intercepted (Fig. 2a). Silicic acid concentrations show a sharp increase from 11  $\mu\text{mol l}^{-1}$  to 55  $\mu\text{mol l}^{-1}$  as latitude changes from 58° S to 61° S and temperature decreases from 3.4°C to -1.3°C in correspondence with the second nitrate increase. A silica maximum (20.1  $\mu\text{mol l}^{-1}$ ) is observed in the cold core centred at latitude 57° S and 58° S. Temperature and nutrient concentrations are roughly constant south of the SACCF (60.1° S).

In November the XBT 2 section shows features similar to those in October. Orthophosphates increase more regularly moving southward, while nitrates and silicic acid are strongly correlated with temperature change ( $R=0.89$ ,  $n=66$ ,  $P < 0.001$ ) and ( $R=0.95$ ,  $n=66$ ,  $P < 0.001$ ), respectively. Nitrate concentrations generally increase from 13  $\mu\text{mol l}^{-1}$  to 28–29  $\mu\text{mol l}^{-1}$  along the section, in particular rising from 14.1  $\mu\text{mol l}^{-1}$  at 55.2° S to 23.8  $\mu\text{mol l}^{-1}$  at 56.2° S as temperature quickly drops from 7.5°C to 4.1°C and where the surface expression of the NSAF (55.9° S) is present. Silicic acid values increase regularly up to 58° S, after which abrupt changes in silicate and temperature are observed. Surface salinity in transects 1 and 2 (Fig. 4c) decreases sharply from 34.40 to 33.85 as temperature cools across the SAF near 55.6–56.0° S, is stable between 56.0° S and 59.4° S and then increases to a plateau of ca. 34.15 from 60° S to 65° S. Similar features are evident in transects 3–5 (Fig. 5) between New Zealand and the western Ross Sea, but the sharp silicate and temperature gradients are shifted between 60° S and 64° S as the PF follows the Pacific–Antarctic Ridge. In November (XBT 3) orthophosphate concentrations show a general increase southward, from values of 0.6–0.8  $\mu\text{mol l}^{-1}$  at 54° S to values of 1.8  $\mu\text{mol l}^{-1}$  at 70° S, as temperature changes from 6–7°C to -1.8°C. The  $\text{NO}_2+\text{NO}_3$  and silicic acid concentrations show a more irregular gradient relative to transects 1 and 2. Silicate, in particular, shows a sharp gradient from 18  $\mu\text{mol l}^{-1}$  (60° S) to 72.5  $\mu\text{mol l}^{-1}$  (65.5° S), while temperatures decrease from 1.7°C to -1.6°C. The abrupt change coincides with the surface expression of the PF.

Orthophosphate concentrations show more variability along the section in December (XBT 4), ranging between 0.8  $\mu\text{mol l}^{-1}$  and 2.8  $\mu\text{mol l}^{-1}$ , with a slight increase southward. The  $\text{NO}_2+\text{NO}_3$  concentrations present a

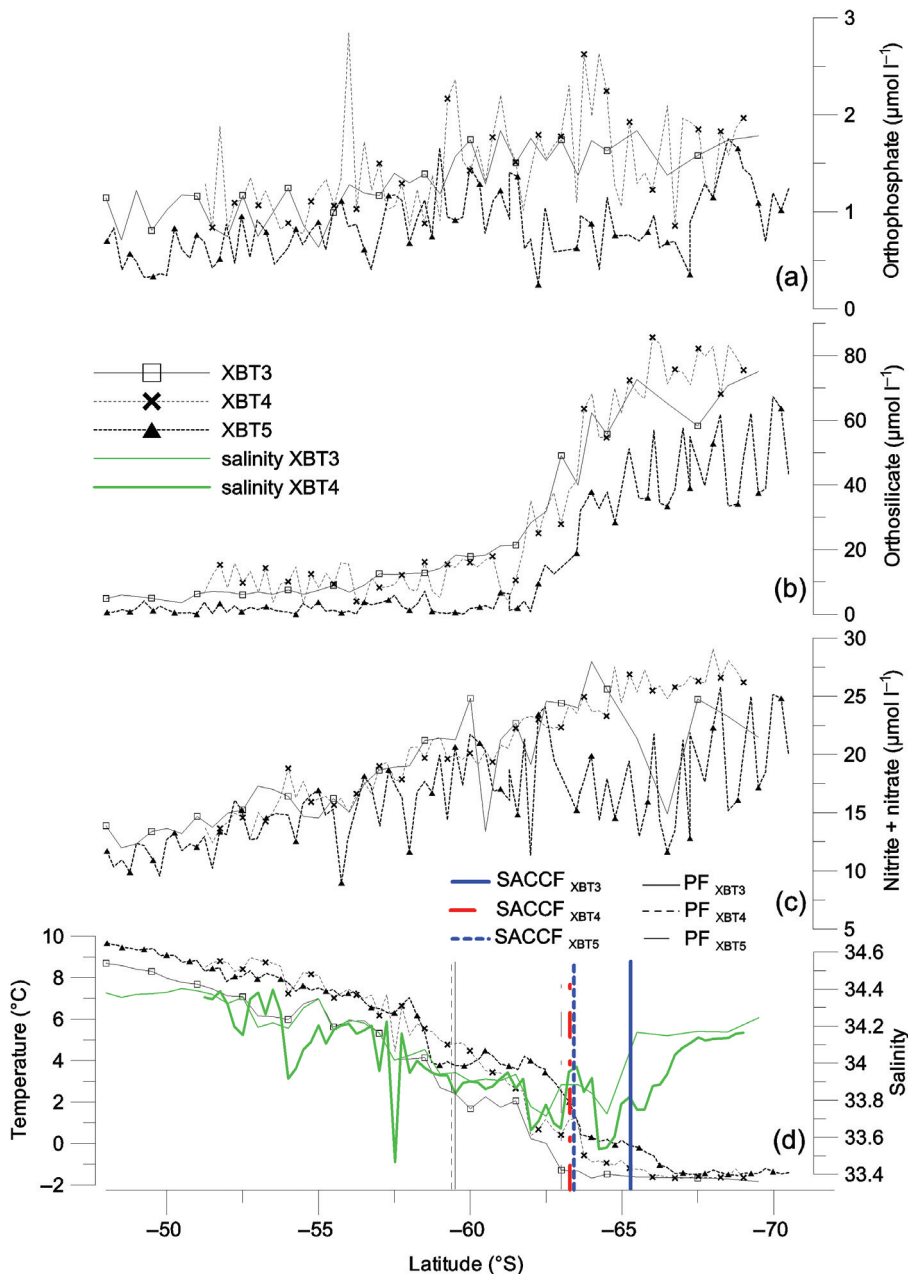
regular increase southward. Silicic acid concentrations increase sharply from 61° S to 64° S as temperature decreases from 3.4°C to -0.9°C, again corresponding to the surface expression of the PF. The salinity gradients in transects 3 and 4 (Fig. 5c) are more regular than in transects 1 and 2, decreasing from 34.35 at 50° S to 33.8 at 60° S, less variable between 60° S and 65° S and increasing to 34.10–34.15 at 67° S and southward, with an overall accompanying temperature decline from +9°C to ca. -1.1°C.

In early January (XBT 5), nutrient concentrations are lower along the entire transect. The  $\text{NO}_2+\text{NO}_3$  and orthophosphate show a more or less regular increase to 62° S (the PF). The  $\text{NO}_2+\text{NO}_3$  range between 10–11  $\mu\text{mol l}^{-1}$  at 48° S and 23–24  $\mu\text{mol l}^{-1}$  at 62° S, while orthophosphates range between 0.4–0.6  $\mu\text{mol l}^{-1}$  and 1.3–1.4  $\mu\text{mol l}^{-1}$  in the same latitude range. South of the PF both nutrients show a slight decrease and then increase regularly. Silicates show different trends, being low between 48° S and 62° S, 0.5–6  $\mu\text{mol l}^{-1}$ . The silicate concentrations then show a sharp increase from 3–4  $\mu\text{mol l}^{-1}$  at 62° S to 30–35  $\mu\text{mol l}^{-1}$  at 65° S as temperature changes from approximately 4°C to 0°C.

## Discussion

Fronts are circumpolar features of the Southern Ocean. On every transect the property fields are found to be organized in such a manner that the transition from warm water in the north to cold water in the south occurs in a series of steps or fronts. These fronts accomplish most of the transport of the ACC and delineate zones characterized by distinct biogeochemical distributions (Sokolov & Rintoul 2007).

It has been shown that there are three distinct circumpolar fronts within the ACC: the SAF, PF and SACCF. The most conspicuous feature in our sections is the SAF, across which temperature at 300 m increases from <4°C to >7°C, south to north. The structure of the SAF is complex with the increase in temperature usually occurring in two steps and the presence of northern and southern branches (see also Russo et al. 1999; Budillon & Rintoul 2003). Locating the front is complicated by possible eddies or meanders associated with these branches, as showed by satellite images of sea surface temperatures (Fig. 6). Following Budillon & Rintoul (2003), we take the northernmost and southernmost expressions of the NSAF and SSAF to define locations of the northern and southern branches. The NSAF can be identified on all XBT sections and is almost always found over the continental slope at the southern edge of Campbell Plateau. This suggests a NSAF slight

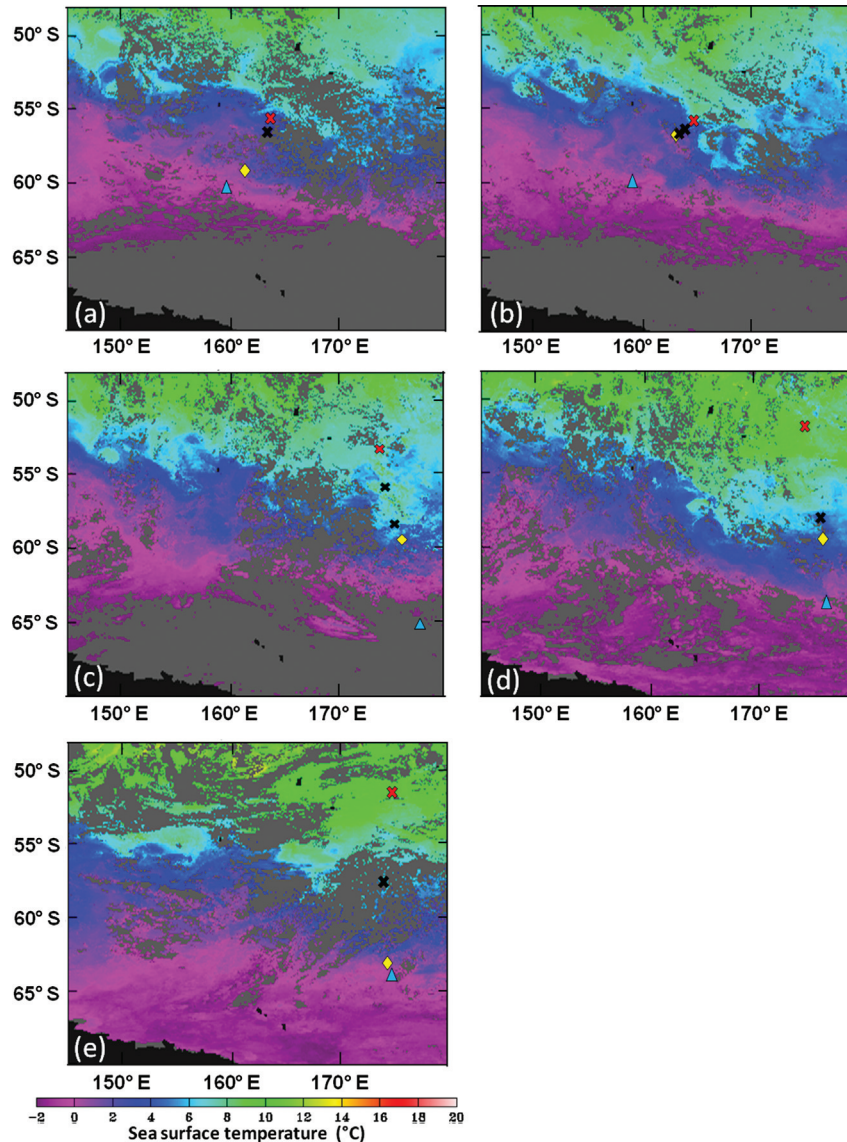


**Fig. 5** Distribution of surface temperature ( $^{\circ}\text{C}$ ), salinity (green line) and nutrients (silicic acid, orthophosphate and nitrite + nitrate [ $\mu\text{M}$ ]) between New Zealand and western Ross Sea from 7 to 15 November 2004 (XBT 3), from 30 November to 5 December 2004 (XBT 4) and from 1 January to 6 January 2005 (XBT 5). (a) Orthophosphate, (b) orthosilicate and (c) nitrite + nitrate. The symbols represent every third sampling event. In the lower panel (d) the vertical lines represent the Polar Front (PF) and the southern Antarctic Circumpolar Current front (SACCF) positions along the transects.

northward shift from November to January along the New Zealand–Ross Sea track. The SSAF shows two expressions in the XBT 2 section, probably due to a meander (Fig. 6b). On the other hand, two expressions of the SSAF along the XBT 3 transect do not appear in the XBT 4 and XBT 5 transects. The position of the PF appears

more variable, with a northward shift in the XBT 2 versus XBT 1 and a southward movement in the XBT 3–5 sections. However, XBT 1 displays a core of cold water with PF characteristics to its north, which should be a meander (Fig. 6a). The SACCF is more stable through time and, to the extent that frontal positions are linked



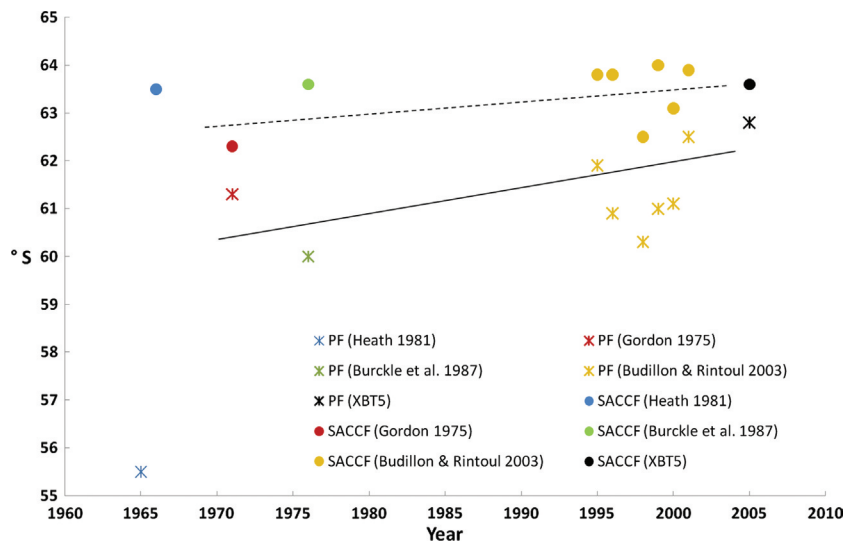


**Fig. 6** Eight-day composites of sea surface temperature ( $^{\circ}\text{C}$ ) from Moderate Resolution Imaging Spectroradiometer (MODIS) images for the study site during the five expendable bathythermograph (XBT) transects: (a) XBT 1, 15–22 October 2004; (b) XBT 2, 31 October–8 November 2004; (c) XBT 3, 8–15 November 2004; (d) XBT 4, 2–9 December 2004 and (e) XBT 5, 1–8 January 2005 (images from <http://poet.jpl.nasa.gov/>). Grey areas represent missing data, primarily due to cloud cover. The positions of the northern Subantarctic Front (red cross), the southern Subantarctic Front (black cross), the Polar Front (yellow diamond) and the southern Antarctic Circumpolar Current front (blue triangle) are indicated.

to the underlying bathymetry (Pollard et al. 2002), their apparent seasonal movements could relate to spatial offsets in the ship tracks.

To evaluate the frontal variability in the studied area through the years, we plotted the positions of the PF and the SACCF found by other researchers working on this topic (Gordon 1975; Heath 1981; Burckle et al. 1987; Budillon & Rintoul 2003) during the same region and season as XBT 5 (Fig. 7). The above authors used different

definitions of the PF and the SACCF so all the positions have been recalculated here using the same definitions. While the Heath (1981) transect along  $160^{\circ}\text{E}$  is less comparable because of the different longitude, the position of the PF moved southward in the period 1965–2005 and the SACCF did not. A southward movement by  $2^{\circ}$  of the PF is quite evident between 1975 and 2005 and may correspond to warming of the Southern Ocean and/or changes in the ACC position during recent



**Fig. 7** Latitude of the Polar Front (PF) and the front of the southern Antarctic Circumpolar Current (SACCF) as observed by Gordon (1975), Heath (1981), Burckle et al. (1987), Budillon & Rintoul (2003) and in our fifth transect. The PF is identified by temperatures below 2°C at 200 m and the SACCF by temperatures below 0°C at depths lower than 150 m and above 1.8°C at depths >500 m (Orsi et al. 1995). The solid line shows the PF regression over the past 40 yrs and the dotted line shows the regression of the SACCF over the same time span; the PF by Heath (1981) was not used for the regression.

decades (Gille 2002). The SACCF shifted less (by 1° of latitude southward), but these changes are comparable to those reported over the Australian–Antarctic Basin (70–150° E) between 1993 and 2006 (Sokolov & Rintoul 2009b). Unfortunately, the relatively poor data from this region that have been published did not permit us to obtain a statistical significant correlation for performing the regression analysis (PF:  $n=9$ ,  $R^2=0.22$ ,  $P > 0.005$ ; SACCF:  $n=9$ ,  $R^2=0.21$ ,  $P > 0.005$ ). More than 85% of the total change in heat content of the Earth’s system has gone into warming the oceans (Levitus et al. 2005). Sokolov & Rintoul (2009b) suggest that the warming of the Southern Ocean is largely caused by a southward shift of the fronts of the ACC. The warming is not uniform around the circumpolar belt as a result of both the spatial distribution of changes in wind forcing relative to the location of the fronts (Sallée et al. 2008) and to interactions between the fronts and bathymetry. By using a 17-year series of remotely sensed sea surface temperatures in the Southern Hemisphere, Kwok & Comiso (2002) demonstrated that positive temperature anomalies and ice-edge retreat in the Pacific sector of the Southern Ocean were associated with large-scale circulation patterns such as El Niño episodes. Our observations indicate that the ACC fronts remained shifted to the south in 2004–05 compared to earlier positions (Gille 2002, 2008; Aoki et al. 2003; Böning et al. 2008; Morrow et al. 2008; Sprintall 2008; Sokolov & Rintoul 2009b), confirming the southward movement of the ACC.

The surface thermal structure plays an important role in the meridional flux of nutrients, as inferred by Budillon & Rintoul (2003). In addition, different surface water masses can be identified by both the temperature and the nutrient trends (Figs. 4, 5), as noted by Morrison et al. (2001). Here  $\text{NO}_2+\text{NO}_3$  and orthosilicate show negative correlations with temperature ( $\text{NO}_2+\text{NO}_3$ :  $n=379$ ,  $R^2=0.69$ ,  $P < 0.001$ ;  $\text{Si}(\text{OH})_4$ :  $n=380$ ,  $R^2=0.77$ ,  $P < 0.001$ ). In the XBT 1 and XBT 2 transects an increase in surface orthophosphate and  $\text{NO}_2+\text{NO}_3$  concentrations is evident in the location of the northern branches of the SAF, as in Rintoul & Bullister (1999). On the XBT 3–5 transects an increase in orthophosphate and  $\text{NO}_2+\text{NO}_3$  concentrations is not evident across the SAF, perhaps due to the increase in light irradiance and the relative increase in nutrient consumption by phytoplankton growth (Mitchell et al. 1991). The surface expression of the PF is well identified by the change in silicic acid concentrations in all the sections, as well as by the Redfield ratio (Table 2). On the XBT 1 transect, silicic acid is high in the cold core between 57° S and 58° S north of the PF as noted above. The  $\text{NO}_2+\text{NO}_3$  also increases in the PF zone but less rapidly than silicate. The  $\text{NO}_2+\text{NO}_3$  and orthophosphate concentrations increase southward but do not show clear correlations with the frontal positions. These nutrient increases are consistent with the presence of high-nutrient, low-chlorophyll water south of the PF, as is the Si/N Redfield ratio (Table 2). Moreover, during the winter period of ice cover, Rubin

**Table 2** Ratio between silicic acid (Si), nitrite+nitrate (N) and orthophosphate (P) in the five expendable bathythermograph (XBT) transects in the different areas: from north to the first expression of the Subantarctic Front (SAF), from the SAF to the Polar Front (PF) and southward of the PF.

	XBT 1	XBT 2	XBT 3	XBT 4	XBT 5
Si/N					
North to SAF	0.45 ± 0.08	0.39 ± 0.05	0.39 ± 0.10	0.69 ± 0.29	0.12 ± 0.09
SAF to PF	0.55 ± 0.19	0.56 ± 0.19	0.55 ± 0.14	0.60 ± 0.26	0.20 ± 0.23
PF to south	1.99 ± 0.37	1.70 ± 0.55	2.10 ± 1.07	2.47 ± 0.63	2.38 ± 0.32
N/P					
North to SAF	15.03 ± 4.34	15.65 ± 6.79	14.56 ± 4.09	14.94 ± 4.58	23.09 ± 7.67
SAF to PF	14.80 ± 3.39	15.80 ± 3.39	16.25 ± 3.37	14.46 ± 4.73	21.45 ± 14.09
PF to south	14.58 ± 2.45	20.43 ± 8.46	13.69 ± 2.31	16.06 ± 5.07	20.93 ± 7.42

et al. (1998) observed in the Ross Sea a low vertical stability that may increase the flux of macronutrients to the surface layer by enhanced deepwater upwelling.

Southward increases in diatom abundance in the Subantarctic region have been reported by several workers in the past (Marumo 1957; Fukase & El Sayed 1965; Burckle et al. 1987; Eynaud et al. 1999). In comparison to XBT 3 and XBT 4, our XBT 5 transect reveals a significant decrease of  $\text{NO}_2 + \text{NO}_3$ , phosphate and silicic acid between  $63^\circ\text{S}$  and  $70^\circ\text{S}$ , as also observed by Morrison et al. (2001) in the same period. High nutrient concentration variability across the XBT 5 section may be due to sea-ice melting. South of  $63^\circ\text{S}$ , orthosilicates average  $42.7 \pm 12.3 \mu\text{mol l}^{-1}$ ; phosphates  $1 \pm 0.3 \mu\text{mol l}^{-1}$  and  $\text{NO}_2 + \text{NO}_3$   $18.2 \pm 4 \mu\text{mol l}^{-1}$ , corresponding to potential nutrient uptakes of  $25 \mu\text{mol l}^{-1}$  (37%),  $0.7 \mu\text{mol l}^{-1}$  (44%) and  $7.4 \mu\text{mol l}^{-1}$  (29%), respectively. These depletions are probably associated with the seasonal phytoplankton bloom characteristic of this period (Moore & Abbot 2002; Green & Sambrotto 2006; Sokolov & Rintoul 2007). SeaWiFS Chlorophyll-*a* concentration images of the five periods evidenced low concentration of Chl-*a* southward of  $50^\circ\text{S}$  between mid-October and mid-November (Fig. 8a–c), while an increase is observed during XBT 4 and 5. Moreover, a phytoplankton increase south of  $63^\circ\text{S}$  can be seen during XBT 5 (Fig. 8e) with respect to the XBT 4 (Fig. 8d) and it is consistent with the observed nutrient depletion and the ice melting typical of the summer season. The use of satellite imagery has drastically changed our view of phytoplankton dynamics. Unfortunately, information gaps caused by cloud and sea-ice cover show the limitations of using remotely sensed data. Nevertheless, Sokolov (2008) demonstrated that within the seasonal sea-ice zone and poleward of the southern branch of the SACCF in particular, the summer chlorophyll blooms are primarily driven by melting of the sea ice, causing formation of very shallow mixed layers, which improves light levels and the release of iron

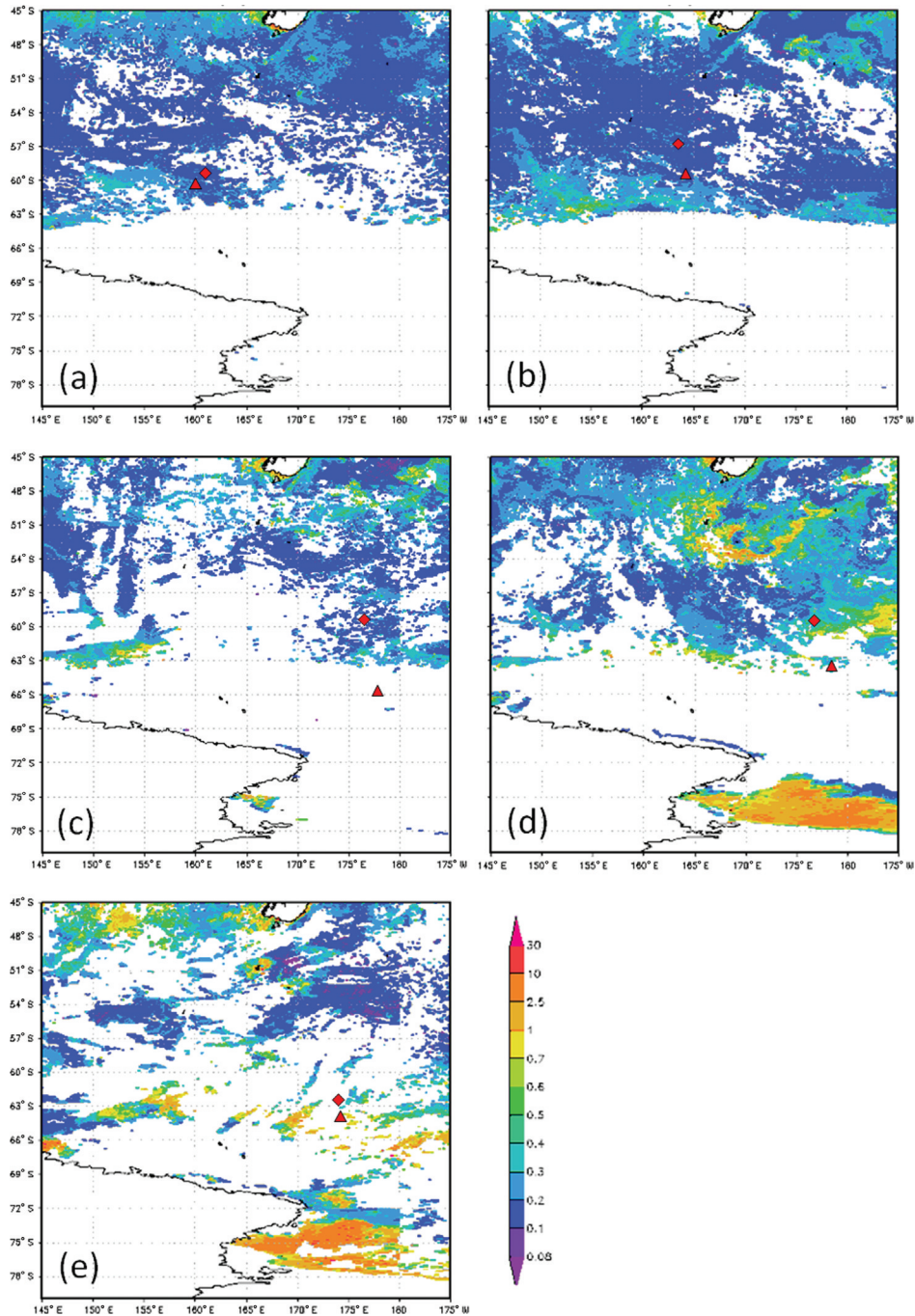
deposited on the sea ice by aeolian deposition during the winter season.

Ratios of apparent nutrient drawdown south of the PF are calculated for these transects in order to estimate bloom and regeneration processes. During the early summer the apparent N/P drawdown ratio is  $8.8 \pm 4.1$  (average  $\pm$  standard deviation) and the apparent Si/N drawdown ratio is  $3.9 \pm 2.8$ . Both values are in accord with a diatom-dominated bloom, as in Sweeney et al. (2000), Morrison et al. (2001) and Green & Sambrotto (2006). Figure 9 shows diatom concentrations derived from the NOBM for all the periods. The model application confirms a diatom growth south of  $63^\circ\text{S}$  that is particularly evident in January 2005 (Fig. 9e).

## Conclusions

The ACC is the major current in the Southern Ocean, isolating the warm stratified subtropical waters from the more homogeneous cold polar waters. Nutrient and XBT observations can be helpful in understanding possible links between thermal and chemical changes of its near-surface fronts and zones. A strong latitudinal gradient in orthosilicate concentration is centred near the PF but also extends well south of that feature as a result of the upwelling of the CDW. Orthophosphate and  $\text{NO}_2 + \text{NO}_3$  concentrations reveal the position of the SAF better than other nutrients during spring and better in spring than summer when phytoplankton growth causes the depletion of these macronutrients. This feature could be due to uptake by different phytoplankton groups not dominated by diatoms or silicoflagellates as well as to the absence of strong orthosilicate input at the surface, as happens south of PF.

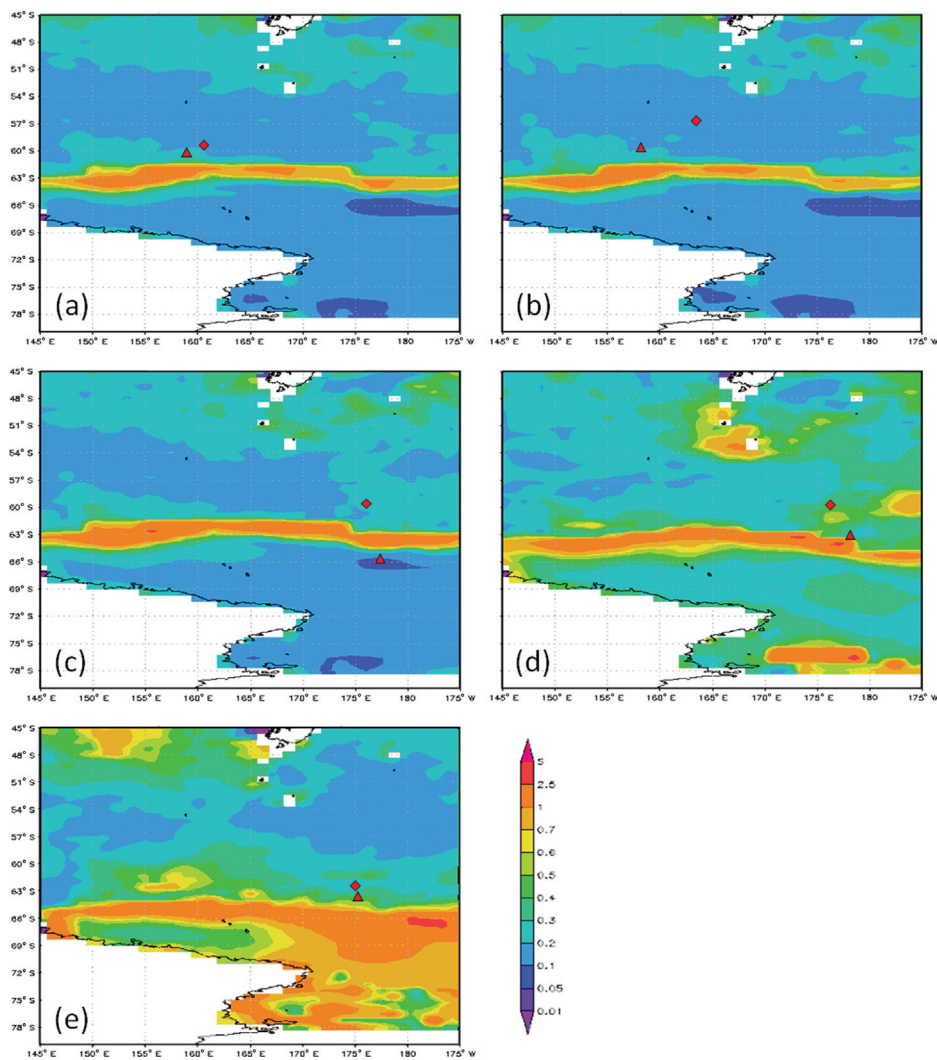
Differences in the density of water masses probably lead to physical obstacles that limit chemical and biological mixing through the fronts within the ACC. During the seasonal phytoplankton bloom, nutrients are significantly depleted south of the PF. The N/P and Si/N



**Fig. 8** Eight-day composites of the National Aeronautics and Space Administration's Sea-viewing Wide Field-of-view Sensor (SeaWiFS) project imagery of chl-*a* ( $\text{mg m}^{-3}$ ) for the study site during the five expendable bathythermograph (XBT) transects: (a) XBT 1, 15–22 October 2004; (b) XBT 2, 31 October–7 November 2004; (c) XBT 3, 8–15 November 2004; (d) XBT 4, 2–9 December 2004 and (e) XBT 5, 1–8 January 2005 (images from <http://reason.gsfc.nasa.gov/Giovanni/>). White areas represent missing data, primarily due to cloud and sea-ice cover. The positions of the Polar Front (red diamond) and the southern Antarctic Circumpolar Current front (red triangle) are indicated.

nutrient drawdown ratios there suggest uptake by diatoms. While the majority of changes in surface nutrient distributions appear to be due to seasonal biological

productivity, exact repeat sections along the same meridian would help to reduce uncertainties introduced by spatial variability.



**Fig. 9** Eight-day composites of the National Aeronautics and Space Administration's Ocean Biogeochemical Model (NOBM) imagery of diatoms ( $\text{mg m}^{-3}$ ) for the study site during the five expendable bathythermograph (XBT) transects: (a) XBT 1, 15–22 October 2004; (b) XBT 2, 31 October–7 November 2004; (c) XBT 3, 8–15 November 2004; (d) XBT 4, 2–9 December and (e) XBT 5, 1–8 January 2005 (images from <http://reason.gsfc.nasa.gov/Giovanni/>). The positions of the Polar Front (red diamond) and the southern Antarctic Circumpolar Current front (red triangle) are indicated.

The positions of ocean fronts in the upper ocean between New Zealand and the Ross Sea in spring–summer 2004–05 were within the ranges that have been observed by other authors (Budillon & Rintoul 2003). Over recent decades, however, the January positions of the PF and SACC front reveal a southward movement of these fronts, perhaps related to ocean warming and a more southward location of the ACC. In particular, the southward movements of the PF positions during 1975–2005 appear evident because in the summer period the ice melt is dominant and more dependent on ocean warming. Since the thermohaline properties of these fronts set the density field and are closely correlated with nutrient levels and

gradients, we infer there have likely been associated ocean current and ecosystem changes.

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