Geothermal resources research in a granitic basement - the Braga area case study (NW Portugal)

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Geothermal resources are increasingly being considered as a strategic alternative in energy production, especially with the latest geopolitical developments. The densely populated Braga region, in NW Portugal, is endowed with a geostructural setting that enables the existence of several thermal water occurrences, spatially associated with a deep-rooted structure - the Vigo-Régua shear zone, set in a granite context. Given the latest advances in aeothermal energy production, it is possible to predict a mid-to long-term implementation of geothermal energy production in the vicinity of that deep rooted structure. Although strongly encouraging, the exploratory geophysical, geochemical and geological data are still insufficient to deliver a definitive frame of the potential energy associated with the estimated reservoirs. Ongoing work combining gravimetric, radiometric and geochemical data will provide a better understanding of the deeply concealed structures.

Les ressources géothermiques sont de plus en plus considérées comme une alternative stratégique dans la production d'énergie, en particulier compte tenu du contexte géopolitique récent. La région densément peuplée de Braga, au nord-ouest du Portugal, est située dans un contexte granitique spatialement associé à une structure enracinée - la zone de cisaillement Vigo-Régua. Ce contexte géologique et structural a permis le développement de plusieurs occurrences d'eau thermale. Compte tenu des dernières avancées en matière de production d'énergie géothermique, il est possible de prévoir une mise en œuvre à moyen et long terme de la production d'énergie géothermique à proximité de cette structure profonde. Bien que fortement encourageantes, les données géophysiques, géochimiques et géologiques exploratoires sont encore insuffisantes pour fournir un cadre définitif de l'énergie potentielle associée aux réservoirs estimés. Des travaux en cours combinant des données gravimétriques, radiométriques et géochimiques permettront de mieux comprendre les structures profondément enfouies.

Los recursos geotérmicos se consideran cada vez más como una alternativa estratéaica en la producción de energía, especialmente con los últimos desarrollos geopolíticos. La región densamente poblada de Braga, en el noroeste de Portugal, está dotada de un entorno estructural que permite la presencia de aguas termales, asociadas espacialmente con una estructura profunda en un contexto granítico: la zona de falla Viao-Réaua. Dados los últimos avances en la producción de energía geotérmica, es posible predecir una implementación a mediano y largo plazo de la producción de energía geotérmica en las cercanías de esa estructura de raíces profundas. Aún cuando los datos geofísicos, geoquímicos y geológicos exploratorios son muy alentadores, son insuficientes para brindar un marco definitivo del potencial geotérmico asociada con los yacimientos. El trabajo en curso que integra datos gravimétricos, radiométricos y geoquímicos proporcionará una mejor comprensión de las estructuras profundas por explorar.

1. Introduction

eothermal resources are considered a sustainable and environmentally friendly alternative to

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* brunopereira@sinergeo.pt produce energy. The European Federation of Geologists (EFG) states that geothermal energy, both shallow and deep, is part of the answer to meet the issue of efficient renewable energy production [1].

The exploration of geothermal resources combines a set of different methodologies and techniques, as described in [2], in which the use of geophysics has an important specific role. However, there is always a significant degree of uncertainty that creates gaps that need to be filled with consistent geological models [3]. Geophysical methods like gravimetry can be used to obtain useful information regarding the deep geostructural setting of geothermal sites, and therefore its fluid circuits. Also, radiometry can be used to detect radiogenic heat generating rocks, potential sources of underground heat.

The north of Portugal seems to be a suitable area for the exploitation of geothermal resources due to its mean geothermal gradient of 35.5 °C/km [4] and a mean heat flow value of 95 mW m^{-2} , derived from borehole measurements [5].

The Braga region, located in NW Portugal, is a zone with well-defined geostructural indicators (*Figure 1*). These indicators are the presence of a deeprooted structure, the Vigo-Régua shear zone, with its associated secondary structural pattern of faults and fractures; the existence of several thermal water occurrences spatially associated with the main faults (*Figures 2,3*), as well as other structural settings; the definition of strong gravimetric anomalies correlatable with the existing geology; and the indication of significant temperatures derived from geothermometry. The application of chemical geothermometers in the region of Braga has allowed the temperature estimate approach of the geothermal reservoirs that feed the researched thermal water occurrences, as well as the its circulation depth. Several studies have demonstrated the suitability of the application of environmental isotopes in the characterisation of geothermal reservoirs in different zones of the Maciço Antigo in northern Portugal [6-10].

The estimated reservoir temperature for the studied sites allows us to predict the feasibility of using direct heat and to promote the eventual electricity production, considering the Lindal diagram [11]. Good results could also be obtained with other complementary techniques, such as electric resistivity and induced polarisation, in specific settings.

This research intends to upscale the existing knowledge regarding the Enhanced Geothermal System (EGS) potential of a specific region of the Variscan granitic basement, as classified in [12], attempting – although in an exploratory approach – to define functional regional guides for geothermal research and exploration.

According to Nicholson [13], systems with temperatures below 150 °C are considered low-temperature systems. The deduced water temperature of the studied Portuguese thermal occurrences is mainly below this threshold, with some values above it. As stated before, these results need to be addressed with caution.

The studied area is densely populated, with annual energy consumption of 3×10^9 kWh in the broader region and a 7×10^8 kWh in the city surroundings in 2020, which is equivalent to 2.1×10^5 tonnes of oil equivalent [14]. The main motivation of this work is to better understand the local deep geostructural setting within the scope of geothermal resources research tools development.

2. Geological setting

The Braga area (Iberian Peninsula; NW Portugal) is located in the Central Iberian Zone (CIZ) of Portugal and in the Galicia–Trás-os-Montes Zone (TMZG) (*Figure 1*) [15].

The CIZ was formed in the Variscan Orogeny, resulting from the continental



Figure 1: Geological-structural setting of the study area.

collision between the continents Gondwana and Laurussia, which began at the end of the Silurian and beginning of the Devonian, characterized by subduction and obduction mechanisms of the oceanic crust [16]. The Variscan cyclebegan with the opening of oceans bordered by passive margins (540-420 Ma), followed by the beginning of the subduction in the Palaeozoic Oceans with the subordinate opening of marginal basins post-arc and ophiolitic blade obduction (420-390 Ma) and high pressure thermometamorphic events. Subsequently, continental collision and orogeny occurred with sedimentary and tectonic polarity oriented towards the foreland zones (390-300 Ma), accompanied by thermal anomalies, generating abundant granitoids and high-temperature metamorphism, followed by a transcurrent intracontinental deformation and localised orogenic collapse (300-270 Ma) [17]. Variscan deformation is characterised by polyphasic processes and is divided into three tectonic phases, D1, D2 and D3 [16,17,18]. The D1 phase generated folds with NW-SE predominant orientation, but in allochthonous, parautochthonous

or autochthonous stratigraphical formations, generates folds with different orientations and vergence, with folds with vertical axial plane in the autochthonous and slightly vergent in the parautochthonous. The D2 phase occurred close to the D1 phase, with the formation of lying folds facing SE being well represented in the allochthonous and the parautochthonous [16], with the development of subhorizontal foliation (S2) [18]. The D3 phase, unlike the previous ones, occurred in all terrains, autochthonous, parautochthonous and allochthonous.

At a regional level, vertical and NNE-SSW oriented corridors and ductile-brittle and brittle shear zones are defined. Still in this phase, previous structures that possibly developed in D1 or D2 (Vigo-Régua shear zone) were reactivated in a transcurrent regime [19]. The Variscan orogeny originated a crustal thickening that produced granitic magmas by anatexis. The Portuguese Variscan granites are synorogenic and can be divided into two groups, two-mica granites are syntectonic relatively to the D3 (syn-D3)



Figure 2: Identified thermal occurrences in the vicinity of the Vigo-Régua shear zone from the Braga region.

Table 1: Data from thermal wate	r occurrences in the Braaa reaion.
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	Temperature (°C)	pН	Depth (m)	Water Type	Use
Caldelas ^(a)	31	8.2	179	Bicarbonate	Spa
Cavadinho/Crespos ^(b)	20.4	9.25	0	Sulphureous	None
Verim ^(b)	<20	8.86	0	Sulphureous	None
Gestal ^(b)	15.5	8.9	0	Sulphureous	None
Caldas das Taipas ^(b)	30	8.0	n.d.	Sulphureous	Spa
Caldas de Moimenta ^(a)	21	9.1	n.d.	Sulphureous	Spa
Caldas de Vizela ^{(a)*}	43.5	9.34	62	Sulphureous	Spa
Gualtar ^(c)	18.0	7.41	0	Sulphureous	None

Legend: (a) [24], (b)[27], (c)[26]; *(b)65°C registered historically

Variscan deformation phase, and are generally located along the core of regional bends D3, with NW-SE direction (D3, last phase of Variscan ductile deformation). They are usually leucocratic granites, with muscovite and primary biotite, resulting from the wet crystallisation of peraluminous magmas, originating from a mesocrustal level [20]. The second group, biotitic granites, originate deeper in the crust and correspond to relatively dried magmas, and if muscovite occurs in this group, it is of secondary origin [21]. Biotitic granites are mainly controlled by D3 shears and late Variscan tectonic structures. The emplacement period of granites, relative to the D3 phase, can be syn-D3 (320-313 Ma), late-D3 (311-306 Ma), late to post-D3 (300 Ma) or post-D3 (299-290 Ma) [22].

In the study area, the relief morphology is marked by tectonics, where the lithologies have been strongly affected by the Variscan NNE-SSW and NW-SE fracturing, and reactivated by Alpine movements that generated their own ENE-WSW fracture network [23]. The granitoids emplacement that outcrops in this area was mainly conditioned by the Variscan phase. Late-Variscan tectonics is marked by the effect of maximum compression with NE-SW orientation caused by NW-SE and fracturing well marked in this area by the great alignments of the river network [23].

The regional setting of the study area can be observed in Figure 1, where the major tectonic-stratigraphic units are identified.

The Vigo-Régua ductile shear zone corresponds to the southernmost segment of a larger system (the ductile shear zone of Malpica-Lamego); (*Figure 1*), with a dominant NW-SE orientation parallel to the Variscan Chain on the NW of the Iberian Peninsula and with a total extension of 275 km [20].

The Vigo-Régua ductile shear zone seg-

ment has a multiphase kinematic interpretation within the Variscan orogeny: left-hand movement in D1 and D2 (370– 310/315 Ma) and right-hand movement in D3 (310/315–300 Ma) [20]. In D1, a thrusting event is identified, followed by an episode changing the structural vergence, previous to D3. This shear zone has a sub-vertical foliation or W-slanted penetration and a sub-horizontal stretching lineation [19]. This shear develops along parautochthonous and associated granitic rocks. The granodiorite rocks are structurally controlled by this shear zone [19].

2.1. The Braga Area

The studied area is located within the Cávado and Ave River watersheds, in NW Portugal (*Figure 2*), included in a broader unit defined as Portuguese Hydrographic Region 2 (RH2). The combined watershed area is of roughly 3,060 km².

The average annual flow on the Cávado and Ave Rivers watershed is about 3,402 hm³. It is estimated that the rainfall in the combined catchment area is on average 1,788 mm/year, ranging from 968 mm to 3,253 mm [23], which guarantees the constant recharge of groundwater systems. There are different thermal water occurrences located near and spatially related to the Vigo-Regua shear zone (*Figure 2*). Data from occurrences surrounding the study area can be observed in Table 1.

The geomorphology of the area is mainly characterised by a typically alveolar morphology with extensive and wide valleys as well as orographic systems with steep slopes, with graben type and reliefs of the Horst as a result of a base geologically markedly granitic and strongly controlled by tectonics [24].

The NNE–SSW and ENE-WSW structures are well marked in the orientation of the rivers and tributaries (*Figure 3*) as well as in the main elevations as a result of the late-Variscan tectonic action and the Alpine movements.

The Cávado deposits occur mainly in the river's right margin, in a graben morphological depression, as well as in the Ave River. The strong NW-SE tectonic control can be observed in the granitic plutons outcropping shapes (*Figure 3*). Also, the E-W valleys tend to be filled with detrital materials, suggesting graben-like structures associated with these late-Alpine valley directions. The metasediments only outcrop on the west block of the Vigo-Régua shear zone, which suggests the upward movement of the East block



Figure 3: Geological map of the study area [25][26][28][29][30].

Table 2:	Water	temper	rature	$(^{\circ}C)$

	Qz geothermom- eter	Na/K geothermom- eter	Measured Tempera- ture
Caldelas(a)	71	153	31
Cavadinho/ Crespos(b)	77	59	20.4
Gualtar	no data	no data	18

Legend: (a)[24] (b) [27]

of this right-hand structure. In the same figure a simplified view of the brittle structures (faults, fissures and gashes) is shown, highlighting crosscutting nodes near known occurrences. Also, the outcropping lithologies are identified.

When observing this figure, it is quite clear that cross-cutting fault and gash structures of WSW-ENE with NNW-SSW directions are associated with the known occurrences, with the exception of the Cavadinho thermal occurrence. This structural criterion was a base for the selection of the three studied occurrences, because of the contrast with the surrounding rock, induced by negative Bouguer anomalies expected in deeply fractured zones [31]. Also, cumulatively, these occurrences are located near urban centres and geothermometric data are available. Isotopic studies [24] and [27] allowed an approach towards the deeper circulation water temperature. The results can be viewed with "optimistic" and "pessimistic" approaches, and need to be addressed with caution. The range of values for water temperature estimates range from 59 °C to 132 °C applying the quartz geothermometer [24] and [27] and from 59 °C to 153 °C using the Na/K geothermometer [24] and [27]. These estimates are well above the natural spring water temperatures, as shown in Table 2. The geology of each one of the three selected sites can be observed

in *Figure 4* (From north to south: Caldelas, Cavadinho, Gualtar).

The Caldelas site geology consists of three outcropping granites, with NW-SE direction, cross-cut by NW-SE and WSW-WNE faults and gashes. The Cavadinho site has only one outcropping granite, with crosscutting faults and gashes with NNE-SSW and WSW-ENE directions. The Gualtar site has a more complex geology with three outcropping granites, as well as metasediment, and is crossed by the Vigo-Regua shear zone.

3. Materials and Methods

The present study is comprised of several tasks: site selection based on crosscutting structures and thermal occurrence locations, collection and interpretation of gravimetric data and interpretation of radiometric data.

An exploratory approach is considered using combined geophysical tools. Considering the extensive demography and highly urban areas on the studied region, the use of electromagnetic and

electric geophysical methods is considered unsuitable due to logistics and signal interference. The choice of geophysical tools fell on the use of gravimetry and radiometry, with the first one aiming to detect density contrasts related to outcropping structures/lithologies that could point out possible reservoirs in low density areas. Radiometry is used with the main objective of revealing radiation anomalies related to highly radiogenic granites [31], which are also structurally controlled. The database [33] was used to model the radiometric background of the study area and selected sites. The gravimetric surveys were conducted using a



Cavadinho

Pioneer, Worden 679, model 155 gravimeter with full DGPS follow-up. Specifics on the conducted gravity surveys can be consulted in References [34] and [35].

4. Results

Figure 5 presents and compares the results from the gravity survey, expressed in complete Bouguer anomaly (mGal) and the radiometric background (nGy/h). A clear relation between radiometric and gravimetric highs can be observed for the Caldelas site. The known thermal occurrence is located in a sudden transition zone between high and low anomalies.

The results obtained for the Cavadinho site (Figure 6) also show that a clear relation between radiometric and gravimetric highs can be observed, although this is not quite as conspicuous as at the Caldelas site. Here also the known thermal occurrence is located in a transition zone between high and low anomalies.

The Gualtar site results (Figure 7) show a more diffuse spatial relation between



Figure 4: Geological setting of the selected sites.



Figure 5: Bouguer anomaly and radiometric background at the Caldelas site.



radiometric and gravimetric highs. The known thermal occurrence is located near a sudden transition zone between high and low anomalies

A broader view of the results can be attained in Figure 8. There is a clear association between known thermal occurrences and radiometric highs.

The comparison of gravimetric and radiometric data points to the existence of a spatial relationship between the location of thermal water occurrences and radiometric maximum values. Similarly, there is an overlap of the gravimetric anomaly maximum values - indicative of higher density zones - and the radiometric maximum values. Thermal water occurrences are also associated spatially with contrasting zones of high and low gravimetric anomalies. Low radiometric values are associated with granodiorites, leucogranites and metasediments from the area, while high values are mainly related with biotite monzogranites.

5. Discussion

The results generated by the conducted exploratory research, although encouraging, must be viewed with caution. In fact, there are many questions that arise from the obtained results. The scale and detail of the gravimetric and radiometric surveys must be increased in order to show small anomalies related to the outcropping geo-structural data. Also, the scale and definition of the geo-structural mapping must be improved, at least in the already studied sites and in the areas near the thermal water occurrences. The now defined regional exploration guides, which consist of overlapping gravitic lows with radiometric highs and cross-cutting brittle structures in the vicinity of a large polyphasic shear zone, must be thoroughly tested with a refining of geophysical surveys, as well as geological mapping combined with seismic data analysis. Despite the high annual water recharge

of the studied area, there is no consensus concerning the origin of thermal water (e.g. [6], [10], [27]), and our results neither support nor reject the hypothesis of deep-water circulation through tectonic structures, which is the most scientifically accepted. The low gravimetric anomalies surrounding the thermal water occurrences are interpreted as the result of the occurrence of porous/fractured rocks acting as a deep reservoir with a not fully understood origin. The highest radiometric values are probably related to highly radiogenic rocks, a possible source of heat [36]. Thus, these occurrences can be the result of the decay of dense radiogenic rocks with nearby reservoirs induced by the tectonic setting endowed by the Vigo-Régua shear zone and associated faults and gashes. The low radiogenic rocks can act as a cap to the heat generated by the most radiogenic ones.

The temperature range of the selected water occurrences is sufficient to allow

nGy/h



Figure 6: Bouguer anomaly and radiometric background at the Cavadinho site.



Figure 7: Bouquer anomaly and radiometric background at the Gualtar site.

electricity generation in binary geothermal power plants [11]. Also, some of the water samples were collected in natural springs and others in boreholes. The comparison of water temperature from boreholes and natural springs allows to state that water temperature increases significantly with depth.

Further studies must be carried out in order to obtain a representative framework of water flow production in these occurrences, which are not known or fully disclosed, mainly the ones that are being currently exploited for balneotherapy and spas.

6. Conclusions

This study was carried out within the Maciço Antigo geotectonic unit in Portugal, focusing on a portion of its Variscan granitic basement which has a geostructural setting that enables the existence of several thermal occurrences. However, the risk of exploration increases considerably in moving away from known hot spring occurrences [36]. There are in-depth geological alteration processes associated with thermal water fluids, unperceived at the surface (non-outcropping), that are detectable through the use of geophysical exploration tools. Local geological singularities will determine the conditions for groundwater circulation and storage. Some parts of the Portuguese territory, namely within its granitic basement, present favourable geostructural indicators for the implementation of stimulated geothermal energy, when compared with other European sites where this technology is being used, despite the large gaps in the detailed knowledge of the underground reservoirs [37].

It is possible to foresee geothermal resource exploration in the vicinities of the deeply rooted Vigo-Régua shear zone, within the Cávado watershed [38], considering the existing data, especially data obtained from geothermometers. However, information on the deep geometry of the Vigo-Régua shear zone and the geothermal fluid circulation mechanisms is scarce. Further research and data collection from combined geostructural, geophysical surveys, hydrogeochemistry and geothermometry of the known regions will allow the determination of drilling targets away from already identified thermal water occurrences and the discovery of geostructural settings favourable for EGS and/or deep geothermal exploration It is also conspicuous that the

estimated temperature calculated from geothermometers and from boreholes is much higher than the water temperature measured at the surface (in springs). This situation needs to be clearly understood, especially considering the deep reservoir and its link to the deeply rooted Vigo-Régua shear zone, assumed to be the structure that controls the heat pulses and geothermal circuits by seismic pumping at a regional scale. Also, this will help the identifcation of potentially anomalous geothermal gradients. The tectonic control of thermal water occurrences is evident, as shown by the radiometric and gravimetric anomalies and their relationship to the configuration of regional structures. These indicators, regional structures, and gravimetric and radiometric anomalies may be used as future guides for the exploration of regional geothermal resources.

The potential of the studied area for Enhanced Geothermal Systems (EGS) and/or deep geothermal is plausible, considering the geological setting, thus enabling the production of heat and eventually electricity from this renewable energy source.



Figure 8: Comparison between radiometric and geological background from the study area and study sites.

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