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Feeding selectivity and rapid particle processing by deep-sea megafaunal deposit feeders: A ²³⁴Th tracer approach

by Robert J. Miller^{1,2}, Craig R. Smith¹, David J. DeMaster³ and William L. Fornes³

ABSTRACT

Deposit-feeding megafauna occur in virtually all deep-sea environments, yet their feeding selectivity and particle processing rates are poorly known. Excess²³⁴Th activity is commonly used as a geochemical tracer for recently settled (<100-d old) particles in the quiescent deep sea, but it has rarely been applied to the study of deposit feeders. To explore the selectivity and rates of megafaunal deposit feeding, we compared excess ²³⁴Th activities in the gut contents of deposit feeders from Santa Catalina Basin (SCB) (\sim 1200 m depth) and the Hawaiian slope (\sim 1680 m) to the activity of surface sediments and, in SCB, to material from sediment traps moored ~ 150 m above the seafloor. We also measured concentrations of chlorophyll a and phaeopigments in animal guts and surface sediments to evaluate feeding selectivity. In the SCB, excess 234 Th (234 Th_{xx}) activities in the guts of four species of surface-deposit feeders were 14-17 fold greater than those of the top 5 mm of sediment. Pannychia moseleyi and Scotoplanes globosa, two highly mobile, surface-deposit-fæding elasipodid holothurians, were the most enriched in gut ²³⁴Th_{xs} activity, suggesting that these species fed very selectively on particles settled to the seafloor within the previous ~ 20 d. *Pannychia moseleyi* guts also exhibited 500-fold enrichment of chlorophyll a relative to surface sediments indicating highly selective ingestion of phytodetritus. Chiridota sp., a burrowing, surface-deposit-fæding, chiridotid holothurian, and Bathybembix bairdii, a surface-deposit-feeding trochid gastropod, were less enriched in gut 234 Th_{xs} activity, reflecting lower mobility and/or less selectivity at time of particle pickup. A subsurface-deposit-feeding, molpadiid holothurian was not enriched in gut ²³⁴Th_{xs} activity compared to surface sediments, but was greatly enriched compared to average activities at its presumed feeding depth of 6-7 cm. On the Hawaiian slope, gut contents of two surface-deposit feeders, the synallactid holothurians Mesothuria carnosa and Paleopatides retifer, were not enriched in ²³⁴Th_{xs} activity; however, M. carnosa and Phryssocystis sp. (a surface-deposit-feeding echinoid) were enriched in chlorophyll a, suggesting that the Hawaiian slope species are also selective feeders. Presumably, frequent sediment resuspension makes ²³⁴Th_{xs} activity a poor tracer for recently settled, food-rich particles on the Hawaiian slope. Based on a newly developed ²³⁴Th-flux model, we calculate that the three dominant megafaunal, surface-deposit feeders in SCB consumed on average 39–52% (s.e. 13–27%) of the daily flux of ²³⁴Th_{xs} activity to the SCB floor. By chemically altering (e.g., digesting) and redistributing recently settled particulate organic matter, these megafauna are likely to substantially influence carbon diagenesis and food-web structure in this bathyal habitat.

1. Department of Oceanography, SOEST, University of Hawaii, 1000 Pope Road, Honolulu, Hawaii, 96822, U.S.A. *email: csmith@soest.hawaii.edu*

2. *Present address:* Biology Department, University of Massachusetts, 100 Morrissey Blvd., Boston, Massachusetts, 02125, U.S.A.

3. Department of Marine, Earth and Atmospheric Sciences, North Carolina State University, Raleigh, North Carolina, 27695-8208, U.S.A.

1. Introduction

Deposit-feeding megafauna occur in virtually all bathyal and abyssal habitats (e.g., Gage and Tyler, 1991). Because these large, relatively sparse animals are rarely collected in quantitative seafloor samples, their effects on community-level processes and sediment geochemistry are easily overlooked. However, time-lapse photographic studies at a number of sites on the continental margin suggest that megafauna may traverse substantial portions of the seafloor over time scales of days to months (e.g., Lafond, 1967; C. Smith, 1992; K. Smith *et al.*, 1993). If these animals selectively ingest recently settled particles, they could process a substantial proportion of the labile particulate-organic-carbon (POC) flux reaching the bathyal seafloor. Such ingestion will physically alter settled particles through gut compaction and fecal-pellet formation (e.g., Taghon et al., 1984), it will change the chemical composition and microflora of fresh particles through digestion and redox alterations (e.g., Plante and Jumars, 1992), and it will concentrate labile POC into fecal deposits and burrow lumens (e.g., Jumars and Wheatcroft, 1989; C. Smith et al., 1986, 1996). In addition, such selective ingestion may lead to age-dependent mixing, wherein recently settled, food-rich particles are mixed faster into the sediment column than are older, food-poor particles (C. Smith et al., 1993). All of these processes will affect the quality and availability of labile POC to other members of the bathyal benthos and will modify the nature and distribution of organic-carbon diagenesis in deep-sea sediments.

The particle selectivity of deep-sea deposit feeders is relatively poorly known, although it is expected to be high. Shallow-water deposit feeders of many taxa, especially those feeding near the sediment-water interface, select particles based on small size, low specific gravity and presence of organic coatings (see reviews in Wheatcroft (1992) and C. Smith *et al.* (1993)). Because the deep sea is a relatively organic-poor environment, with labile POC primarily sinking from overlying waters, deep-sea deposit feeders are likely to preferentially consume recently settled particles using mechanisms similar to their shallow-water counterparts (C. Smith *et al.*, 1993). In fact, a variety of megafaunal deposit feeders from the abyss are known to selectively ingest or sequester chlorophyll-*a*-rich, phytodetrital material that is newly arrived at the seafloor (Billet *et al.*, 1988; Thiel *et al.*, 1989; Moore and Roberts, 1994; Roberts *et al.*, 1996; C. Smith *et al.*, 1996). Diatom-addition experiments using ¹³C-labeled algae also imply rapid uptake of fresh phytodetritus by a variety of macrofaunal surface-deposit feeders from bathyal depths (Levin *et al.*, 1997).

In this study we use excess 234 Th (234 Th_{xs}) activity as a natural tracer to explore particle selectivity and particle processing rates of bathyal deposit feeders. 234 Th_{xs} is scavenged by particles settling through the water column and has a radioactive half life of 24.1 d; it thus may be used to trace the behavior of particles for roughly 100 d after they have reached the seafloor (e.g., Aller and DeMaster, 1984; C. Smith *et al.*, 1993). In earlier work, we predicted that in quiescent deep-sea habitats, deposit-feeder gut sediments should be substantially enriched in excess 234 Th activity due to selective feeding on recently settled, food-rich particles (C. Smith *et al.*, 1993). A recent study (Lauerman *et al.*, 1997) has, in



Figure 1. Locations of study areas on the Hawaiian slope (circle) and in Santa Catalina Basin (triangle). Contours are in meters.

fact, found high excess ²³⁴Th activities in the gut sediments of two abyssal species of holothurians on the California margin.

Here we further test our predictions of feeding selectivity in megafaunal deposit feeders using ²³⁴Th_{xs} and chlorophyll-*a* measurements in the quiescent Santa Catalina Basin, and on the physically more energetic Hawaiian slope. In particular, we compare the feeding selectivity of eight species of deposit feeders from three classes (holothurians, echinoids and gastropods) belonging to the three different feeding guilds (highly mobile, epibenthic surface-deposit-feeders; borrowing surface-deposit feeders; head-down subsurface-deposit feeders). In Santa Catalina Basin, we also develop a new tracer approach, using excess ²³⁴Th activity, to evaluate the processing rates of recently settled particles by megafaunal populations. We find that the often overlooked deposit-feeding megafauna can feed very selectively on recently settled (or "young") particles and could substantially influence the diagenesis and availability of labile POC in at least some bathyal and abyssal habitats.

2. Study sites

Santa Catalina Basin (SCB) is a flat-bottomed basin of about 1220 m depth (Fig. 1); our study area was a square roughly 8 km on a side centered at 33° 9.5′ N, 118° 28′ W. Actual stations ranged in depth from 1212 to 1240 m (station details are available upon request from C. Smith). Bottom sediments in SCB are poorly sorted clayey silt, with sedimentation

rates of 12–21 cm per 1000 years (Emery, 1960; DeMaster *et al.*, in prep.). Bottom-water oxygen concentrations are 18–23 micromolar (Archer *et al.*, 1989), the physical environment is relatively low-energy (we have never observed currents of erosive magnitudes in over 50 submersible dives and 3 months of time-lapse camera deployments), and the fauna is well described (C. Smith, 1985, 1986). Deep POC flux in the general area appears to be relatively constant throughout the year [e.g., ranging from 0.02–0.06 g/m²/day in the nearby San Pedro Basin (Thunnel *et al.*, 1994)].

The Hawaiian slope site $(19^{\circ} 36' \text{ N}, 156^{\circ} 4' \text{ W})$ is west of the island of Hawaii, on a ~ 1680 -m deep platform at the base of a 22° slope (Fig. 1). During two submersible dives, we saw no evidence of sediment transport, although sediments were difficult to penetrate with cores. Our sampling area ranged in depth from 1680 to 1685 m and was sediment-covered, with occasional rock outcrops. Hawaiian waters are characterized by relatively low productivity and seasonality (e.g., Karl and Lukas, 1996). The deep-sea megafauna of this region is described in Chave and Jones (1991).

3. Methods

a. Field methods

In SCB, we collected five common species of deposit-feeding megafauna in Dec 1995 and Jan 1996: the epibenthic surface-deposit feeders Pannychia moseleyi, Scotoplanes globosa (elasipodid holothurians), and Bathybembix bairdii (trochid gastropod); the burrowing, surface-deposit-feeding holothurian Chiridota sp.; and the head-down (conveyorbelt) deposit-feeder *Molpadia* sp. (molpadiid holothurian). On the Hawaiian Slope in Oct 1996, we studied the epifaunal surface-deposit feeders Mesothuria carnosa (synallactid holothurian), Paleopatides retifer (synallactid holothurian), and Phryssocystis sp. (echinoid). Epibenthic megafauna were collected in SCB with the submersible *Alvin* (dive nos. 3048-3054) and on the Hawaiian slope using the Hawaiian Undersea Research Laboratory submersible *Pisces V* (dive nos. 317-318) with a scoop net. Megafauna in SCB (in particular Molpadia sp., Scotoplanes, and Chiridota sp.) were also collected with a 8-m semi-balloon otter trawl (outer mesh 3.8 cm, inner mesh 1.3 cm) deployed from the RV New Horizon or RV Atlantis II. Chiridota sp. were also recovered from several USNELtype, 0.25 m² box-core samples. Animals were dissected within two hours of recovery and the gut contents frozen. Material used for pigment analysis was frozen in liquid nitrogen and stored at -80° C. Holothurian foreguts were identified by the restricted area separating them from the "hindgut" [i.e., the intestine, rectum and cloaca (Roberts et al., 1996)]. Depending on the size of the individual gut, hindgut contents were frozen whole or divided into 2-4 sections. Owing to small animal size, we combined gut sediments of Chiridota sp. (10 individuals per group) and Bathybembix bairdii (3-5 per group) to provide adequate mass for radiochemical analyses.

In SCB, sediment cores were collected with *Alvin* using an Ekman box corer (20×20 cm, with 4 10 × 10-cm subcores) or *Alvin* tube corer (6.8-cm diameter). In addition, 10-cm diameter cores were taken in SCB using a multiple corer (Gage and Tyler, 1991). Tube

cores (6.8-cm diameter) were taken on the Hawaiian Slope with the *Pisces V*. Cores were extruded and sectioned into depth intervals as in C. Smith *et al.* (1993), cutting away about 5 mm of the periphery of each depth interval to avoid contamination from vertical smearing. In SCB, cores were sectioned at 0.5-cm intervals to a depth of 2.0 cm, and at 1-cm intervals below 2 cm. On the Hawaiian slope, cores were sectioned into 0–0.2 and 0.2–0.5 cm depth intervals, and then at 0.5 cm intervals to the core bottom. Depth intervals were homogenized, split for analyses and frozen. Pigment samples were frozen in liquid nitrogen and stored at -80° C.

Replicate, single-cup sediment traps of the type used by Dunbar *et al.* (1989) were deployed 150 and 170 m above bottom in SCB for 32.8 days beginning on 5 December 1995. Trap samples were preserved *in situ* using a hypersaline formalin solution (5% NaCl w/w, 10% formalin by volume), and swimmers were removed from samples after recovery.

To evaluate the abundance of epibenthic megafauna in SCB, photographic transects were conducted using *Alvin*'s sponson camera as in Smith and Hamilton (1983), with *Alvin* skiing along the bottom to maintain constant height, traveling at ~15 m/min on a random heading, with an exposure interval of 15 s. The three phototransects were an average of 537 ± 56 (s.e.) m long with a mean of 143 ± 15 photographic frames each.

b. Laboratory methods

 234 Th_{xs} activity was measured using a variation of the technique of Aller and DeMaster (1984), as described in Pope (1992). Briefly, 234 Th activity was isolated and then measured on a low level, gas flow, anti-coincidence beta counter. 238 U activities were measured by alpha spectroscopy. 234 Th activities were converted to excess values by subtracting 238 U activity from total 234 Th activity and correcting for decay since sample collection. All excess 234 Th activities are expressed in dpm g⁻¹ dry sediment (corrected for salt content).

For pigment analyses, samples of frozen gut and core sediments (~ 2 g) were vortexed, sonicated in a dark ice bath for 10 minutes, and extracted overnight at -20° C in the dark in 10 ml of 100% acetone in glass tubes (Stephens *et al.*, 1997). Acetone extraction may reduce degradation artifacts associated with fluorometric measurement of chlorophyll-*a* (Jeffrey and Hallegraeff, 1987). Samples were then centrifuged for 5 min and the absorption of the supernatant at 670 nm read on a Turner model 10-AU fluorometer. Phaeopigments were measured using the acidification technique (Strickland and Parsons, 1972). Chlorophyll *a* (chl-*a*) and phaeopigment concentrations were calculated based on a standard solution from Fisher Scientific. Values were normalized to salt-corrected dry weights of extracted sediment. Sun *et al.* (1991) reported 90.6 ± 3.1% recovery of standard chl-*a* from wet sediments using fluorometry; we judged this level of accuracy acceptable for measuring concentrations of chl-*a* as an indicator of the presence of fresh phytodetritus.

Gut enrichment factors (EF) for ²³⁴Th_{xs} activity and chl-*a* concentrations were calculated as the ratio of gut activity or concentration to surface-sediment activity or concentration. Average gut enrichment factors were calculated for each sampling period (Dec 1995 and Jan 1996) using surface sediment collected at the same time (Table 1). Overall means were

Table 1. ²³⁴Th_{xs} activities, chlorophyll *a* concentrations, mean enrichment factors (EF) and apparent ²³⁴Th_{xs} ages of SCB surface and gut sediments compared to corrected sediment-trap activity. Chlorophyll *a* gut concentrations are for foreguts; enrichment factors for them are as compared to surface sediment. SCB is Santa Catalina Basin, and HS is Hawaiian Slope. Data are means \pm one standard error.

		²³⁴ Th _{xs}			Chl a			Apparent 234Th _{xs}
Sample	Date	(dpm/g)	EF	п	(µg/g)	EF	п	age (d)
SANTA CATALINA BASIN:								
Sediment trap (uncorrected)	12/5/95–1/7/96	1260 ± 95		2				
Sediment trap (corrected)	12/5/95-1/7/96	1960 ± 150		2				0
Surface sediment (0-5 mm)	12/95	30.9 ± 2.4		3				143
Surface sediment (0-5 mm)	1/96	4.9 ± 0.9		3	0.36 ± 0.09		2	207
Subsurface sediment (6-7 cm)	12/95	0 ± 0		3				
Subsurface sediment (6-7 cm)	1/96	0 ± 0		6				
Pannychia moseleyi	12/95	2730 ± 300	88 ± 11	3				-11.4
Pannychia moseleyi	1/96	1290 ± 90	261 ± 57	6	200.77 ± 71.90	558 ± 399	4	14.4
Chiridota sp.	12/95	767 ± 50	25 ± 6	3				32.4
Chiridota sp.	1/96	114 ± 7	23 ± 1	3				98.1
Bathybembix bairdii	12/95	883 ± 90	29 ± 13	3				27.5
Bathybembix bairdii	1/96	144 ± 14	29 ± 6	2				90.0
Scotoplanes globosa	12/95	2290 ± 195	74	1				-5.4
Molpadiid holothurian	12/95	4.1 ± 0.5	0	3				212.7*
HAWAIIAN SLOPE:								
Surface sediment (0-2 mm)	10/96	2.0 ± 0.9		1	0.44 ± 0.01		3	
Phryssocystis sp.	10/96				5.15 ± 0.43	12 ± 4	3	
Mesothuria carnosa	10/96	3.2 ± 0.5	2 ± 1	2	1.18 ± 0.34	3 ± 2	2	
Paleopatides retifer	10/96	2.9 ± 1.1	1 ± 1	3				

*Age compared to surface sediment is 51 d.

calculated as the average of the means from the two sampling period (Dec 95 and Jan 96, yielding n = 2) (Fig. 2), again compared with the mean of sediment from the same time period.

Population densities of *Pannychia moseleyi, Chiridota* sp., and *Bathybembix bairdii* in SCB were estimated from *Alvin* photographic transects. Each photographic frame was projected onto a perspective grid (McNeil, 1954), provided by the *Alvin* Group, using a Besseler model 23CII enlarger. Animals were counted from a 2.25 m² square in the central foreground of each frame where they were clearly identifiable. Mean organism density for each transect was calculated as an independent estimate of abundance for the study area. The three transect means were then averaged to yield an overall mean and variance.

Because *Chiridota* sp. could not be resolved in transect photos due to its small size and burrowing habits, its densities were estimated from seven sets of vertical, close-up, time-lapse photographs taken at 0.8–8 hr intervals from a free-vehicle camera system (the "moundster" camera) deployed at our SCB site in 12/86 (n = 3), 1/87 (n = 2) and 10/87

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Figure 2. Top: Mean ${}^{234}\text{Th}_{xs}$ activities in the surface sediment (0–5 mm), deposit-feeder guts, and sediment-trap material from Santa Catalina Basin. Note that ${}^{234}\text{Th}_{xs}$ activity is plotted on a log scale. Where possible, means given are the average of the two means for the Dec and Jan sampling times; *Scotoplanes, Molpadia* and sediment-trap samples were collected only at one time so the mean of all individuals collected is given. Error bars = one standard error, with the exception of *Scotoplanes globosa* (n = 1), where the error is a 1 σ counting error (see text). Numbers above the histograms are mean enrichment factors relative to surface sediment. Bottom: Mean chlorophyll-*a* and phaeopigment concentrations in surface sediments and *P. moseleyi* guts in Jan 1996. Mean enrichment factors relative to surface sediments are indicated by numbers above bars. Samples sizes are given in parentheses.

(n = 2) (C. Smith *et al.*, 1986; C. Smith, 1992). Each photograph covered an area of 0.3 m², and *Chiridota* sp. were clearly visible as they extended from burrows to feed (C. Smith, 1992). Because this holothurian maintains only one feeding opening at the terminus of its horizontal burrow (C. Smith, 1992), the number of active burrow openings visible in the first frame of each series was used to estimate *Chiridota* sp. density. Active openings were identified by visible extension of a chiridotid in the first, or in a closely subsequent, frame.

c. Flux consumption estimates

We used two approaches to estimate the processing rate of recently settled particles by *P. moseleyi, Chiridota* sp., and *B. bairdii* in SCB: a *thorium-flux model* that we developed, and an *end-member mixing model* similar to that of Lauerman *et al.* (1997). Our

thorium-flux model estimates the flux of 234 Th_{xs} activity through the guts of deposit feeders, using the following equation:

$$TF = \frac{a \cdot gw \cdot D}{gt}$$

where TF is the flux of ²³⁴Th_{xs} activity through the animal population (dpm m² d⁻¹), *a* is the mean weight-specific ²³⁴Th_{xs} activity of the gut contents (dpm g⁻¹), *gw* is the average dry weight of gut sediments for the species in question (*g*), *D* is the population density (m⁻²), and *gt* is the gut throughput time (d). *TF* was compared to the flux of ²³⁴Th_{xs} to the seafloor as measured by the sediment traps and as estimated from seafloor inventories of ²³⁴Th_{xs} (Fornes *et al.*, in prep.). Assuming steady state, the seafloor flux equals the seafloor ²³⁴Th_{xs} inventory (obtained by integrating sediment profiles of ²³⁴Th_{xs}) multiplied by λ , the radioactive decay constant for ²³⁴Th (0.029 d⁻¹).

As a comparison to the thorium-flux model, we also estimated consumption of the "fresh" particulate flux for the same three species in SCB using an *end-member mixing model* (Lauerman *et al.*, 1997). We calculated the proportion, P, of freshly deposited material in the gut of an animal using the following equation (cf. Lauerman *et al.*, 1997):

$$P = \frac{(a - ss)}{(C_o - ss)}$$

where $ss = \text{mean} \, ^{234}\text{Th}_{xs}$ activity of surface sediment (0–5 mm), and $C_o =$ the mean $^{234}\text{Th}_{xs}$ activity (dpm g⁻¹) of particulate material at time of capture in the sediment trap, i.e., "fresh," or zero-age material. This model assumes that the animals are feeding either on zero-age material or bulk surface sediment, that $^{234}\text{Th}_{xs}$ is homogeneously distributed among particles caught in the sediment traps, and that $^{234}\text{Th}_{xs}$ is a conservative tracer relative to gut throughput times (≤ 24 hr, see below). Measured $^{234}\text{Th}_{xs}$ activity of the sediment-trap material was corrected to an average time-of-capture, or zero-age, activity (C_o) using the following equation:

$$F \cdot T \cdot C_{avg} = \int_0^T C_o \cdot F \cdot e^{-\lambda t} dt$$

where C_{avg} is the measured bulk ²³⁴Th_{xs} activity of trap material, *F* is the mass flux (g d⁻¹) into the trap, *T* is trap deployment interval (in this case, 32.8 days), and *t* is time since trap deployment. C_o and *F* are assumed to be constant over the trap deployment. Solving the above equation:

$$C_o = \frac{\lambda \cdot T \cdot C_{avg}}{(1 - e^{-\lambda T})}.$$

Fresh particulate flux, A (dpm m⁻² d⁻¹) through the population was then estimated using the following equation:

$$A = \frac{P \cdot g_W \cdot D}{gt}$$

(1970).

Published throughput times for deep-sea holothurians range from 16–24 hours (Deming and Colwell, 1982; Sibuet, 1988). In situ tracer experiments in SCB with P. moseleyi indicate throughput times of ~12 h (Galley, Smith and DeMaster, in prep.). Thus, we used 12 hours as the throughput time for P. moseleyi and the much smaller Chiridota sp. (cf. Penry and Jumars, 1987). Deposit-feeding gastropods have been found to ingest 0.4 to 25 times their body weight in sediment daily (Lopez and Levinton, 1987). Bathybembix bairdii has an average gut-content weight of $49 \pm 3.3\%$ (\pm s.e.) of dry body weight (unpublished data). If this species ingests 0.4–25 times its body weight per day, this translates to gut passage times of 0.4–20 h. Since deep-sea gastropods are likely to have gut throughput times roughly comparable to their shallow-water counterparts (Penry and Jumars, 1987), we used 24 hours as a conservative estimate of throughput time for B. bairdii. Mean gut dry weights were obtained by dissecting formalin preserved specimens of each species, and drying the guts at ~60°C for >24 hr. Standard errors were propagated through all calculations as relative errors using the methods of Hodges and Lehmann

We also calculated the mass flux of sediment through *P. moseleyi*, *Chiridota* sp., and *B. bairdii* using a *mass-flux model*. Mass flux, M (g m⁻² d⁻¹), through a megafaunal population was estimated as follows:

$$M = \frac{gw \cdot D}{gt}$$

Apparent ²³⁴Th ages of sediments relative to reference material (sediment-trap material or surface sediment) were calculated using the following equation:

$$t_a = \frac{\ln A_0 - \ln A_t}{\lambda}$$

where t_a is apparent age, A_0 is the ²³⁴Th_{xs} activity of the reference material at time zero and A_t is the ²³⁴Th_{xs} activity of the sample (=activity of the reference material at age *t*).

4. Results

a. ²³⁴Thorium measurements

Surface-deposit feeders in Santa Catalina Basin had gut ²³⁴Th_{xs} activities 23–261 times greater than that of average surface (0–5 mm) sediment (Table 1, Fig. 2). The epibenthic holothurians *P. moseleyi* and *S. globosa* had the highest gut ²³⁴Th_{xs} activities, with corresponding apparent gut-sediment ages, relative to sediment-trap material, of -11 to 14 d (Table 1). The gastropod *B. bairdii* and the burrowing holothurian *Chiridota* sp. had lower gut ²³⁴Th_{xs} activities, and apparent gut sediment ages of 28 to 98 d (Table 1). The molpadiid holothurian, a subsurface-deposit feeder, had an average gut ²³⁴Th_{xs} activity lower than that of surface sediment at the time of collection (Dec), with an apparent gut-sediment age relative to surface sediment of 51 d (Table 1, Fig. 2). Nevertheless, the gut activity of the molpadiid was substantially higher (mean = 4.3 dpm g⁻¹ vs 0 dpm g⁻¹)



Figure 3. Mean ${}^{234}\text{Th}_{xs}$ activities (top), and chlorophyll-*a* and phaeopigment concentrations (bottom) of surface sediment (0–2 mm) and deposit-feeder guts on the Hawaiian slope. Note that ${}^{234}\text{Th}_{xs}$ activities are plotted on a log scale. Numbers in parentheses are sample sizes. Mean enrichment factors relative to surface sediments are indicated by numbers above bars. Error bars are one standard error with the exception of the ${}^{234}\text{Th}_{xs}$ for surface sediment (*n* = 1), which is a 1 σ counting error (see text).

than mean sediments at the 6-7 cm depths at which we presume these 6-7 cm long, head-down deposit feeders to be feeding (Table 1).

Gut sediments of two surface-deposit-feeding holothurians from the Hawaiian slope, *M. carnosa* and *P. retifer*, showed no significant enrichment of ²³⁴Th_{xs} activity relative to surface sediment (Table 1, Fig. 3). It should be noted, however, that the surface sediment value is based on a single sample, and seafloor ²³⁴Th_{xs} activities often are highly variable within a site.

The corrected flux of 234 Th_{xs} activity into the sediment trap in SCB for the period 12/5/95-1/7/96 was 731 ± 6 dpm m⁻² d⁻¹, and the mass flux was 0.364 ± 0.02 g m⁻² d⁻¹. Based on seafloor inventories (Fornes, 1999; Fornes *et al.*, in prep.), the flux of 234 Th_{xs} activity to the SCB floor in December, 1995 was 1830 ± 160 dpm m⁻² d⁻¹, while in January, 1996, it was roughly 5-fold lower, i.e., 374 ± 38 dpm m⁻² d⁻¹. Mean surface-sediment and species-specific gut activities of 234 Th_{xs} in SCB were also 2–6 fold higher in December than in January (Table 1).



Figure 4. Mean chlorophyll-*a* and phaeopigment concentrations (± 1 s.e.) along the guts of *Pannychia moseleyi* specimens from SCB, collected Mar 1996. Three measurements were made for each gut section.

b. Chlorophyll-a measurements

Concentrations of chl-*a* and phaeopigments were measured in the foreguts of one species in SCB, *P. moseleyi*. The average chl-*a* concentration was >500 times that of surface sediment (Table 1, Fig. 2). Chlorophyll-*a* was measured along the gut to provide indications of digestion and/or patchiness of food resources. Concentrations generally decreased along the *P. moseleyi* gut, with the foregut having the highest chl-*a* values in three of four cases (Fig. 4). On average, hindgut concentrations of chl-*a* were \sim 38% those in the foregut (Fig. 5). Phaeopigment concentration within the gut followed very similar trends.

On the Hawaiian slope, the foreguts of *M. carnosa* and *Phryssocystis* sp. had average chl-*a* concentrations 2.7 and 12 times greater than surface sediment, respectively (Fig. 3). Phaeopigments were also elevated 1.3–4 fold relative to surface sediments.

c. Population densities and mean gut weights

Based on time-lapse and survey photography, *Chiridota* sp. was by far the most abundant of the three prominent megafaunal surface-deposit feeders that we studied in SCB (Table 2). *Bathybembix bairdii* was approximately an order of magnitude less abundant than *Chiridota* sp., and *P. moseleyi* was roughly two orders of magnitude less abundant (Table 2). The mean individual weights of whole-gut sediments for these species



Figure 5. Mean (± 1 s.e.) percentage of foregut chlorophyll-*a* concentrations in the fore-, mid- and hindguts of *Pannychia moseleyi* from SCB (n = 4). For individuals numbers 1, 2 and 4 (Fig. 4), gut sections 1 and 2 were combined for the midgut data, and sections 3 and 4 for the hindgut data.

were inversely related to population density, yielding roughly equivalent standing stocks of gut sediments within the three populations (Table 2).

d. Flux consumption estimates

Consumption rates of ${}^{234}\text{Th}_{xs}$ fresh particles, and sediment mass by *P. moseleyi*, *Chiridota* sp., and *B. bairdii* are summarized in Table 3. Based on the *thorium-flux model*, the three species combined consumed $380 \pm 200 \text{ dpm d}^{-1}$ of ${}^{234}\text{Th}_{xs}$ activity over Dec–Jan; this was about 50% of the daily sediment-trap flux of ${}^{234}\text{Th}_{xs}$ activity, and about 40% of the

Tabl	e 2.	Mean	individual	gut w	eights,	populati	on dens	sities,	and	standin	g stocks	of	gut se	ediments
(±	-1 st	andaro	d error) for	the the	ree don	ninant me	egafaun	al sur	face	-deposi	t feeders	in S	Santa	Catalina
Ba	asin.													

Species	Gut dry weights (g)	Density (no./m ²)	Gut-sediment standing stock (g/m ²)
Pannychia moseleyi	1.94 ± 0.36 (<i>n</i> = 12)	0.026 ± 0.02	0.05 ± 0.04
Chiridota sp.	0.07 ± 0.02 (<i>n</i> = 30)	2.4 ± 0.6	0.17 ± 0.09
Bathybembix bairdii	0.3 ± 0.03 (<i>n</i> = 10)	0.164 ± 0.12	0.05 ± 0.04

Table 3. Summary of the fluxes of ${}^{234}\text{Th}_{xs}$, "fresh particles" and mass through the guts of animals (i.e., ingested), and the ratios of these gut fluxes to the respective sediment trap and seafloor fluxes. Means \pm one standard error are given. The ${}^{234}\text{Th}_{xs}$ activity of sediment-trap material was 1960 \pm 150, as corrected to average time zero activity (see text), and the flux was 731 \pm 6 dpm/m²/d. The ${}^{234}\text{Th}_{xs}$ flux calculated using seafloor inventories was 1830 \pm 160 dpm/m²/d (Dec.) and 374 \pm 38 dpm/m²/d (Jan.) (Fornes *et al.*, 1999). Particulate (mass) flux, measured via sediment traps, was 0.364 \pm 0.02 g/m²/day.

		Thorium-flux model			End-me	ember mixing r	Mass-flux model		
Species	Date	234 Th _{xs} ingested (dpm m ⁻² d ⁻¹)	% of trap ²³⁴ Th _{xs} flux ingested	% of seafloor ²³⁴ Th _{xs} flux ingested	% of fresh material in gut	Fresh particles ingested (g m ⁻² d ⁻¹)	% of fresh particle flux ingested	Mass ingested (g m ⁻² d ⁻¹)	% of trap mass flux ingested
Pannychia moseleyi	Dec. 1995	280 ± 180	38 ± 24	15 ± 9.8	100 ± 19	0.10 ± 0.08	27 ± 23		
Chiridota sp.		260 ± 100	36 ± 14	14 ± 5.8	38.2 ± 3.9	0.13 ± 0.07	36 ± 19		
Bathybembix bairdii		44 ± 30	6 ± 4	2.4 ± 1.7	44.2 ± 5.8	0.02 ± 0.02	6 ± 5		
Т		580 ± 300	80 ± 42	32 ± 17		$\textbf{0.25} \pm \textbf{0.17}$	69 ± 47		
Pannychia moseleyi	Jan. 1996	130 ± 83	18 ± 11	36 ± 23	65.7 ± 6.8	0.07 ± 0.05	18 ± 15		
Chiridota sp.		38 ± 15	5 ± 2	10 ± 4.1	5.6 ± 0.6	0.02 ± 0.01	5 ± 3		
Bathybembix bairdii		7.5 ± 5.0	1 ± 0.7	2.0 ± 1.4	7.1 ± 0.9	0.00 ± 0.00	1 ± 1		
Т		180 ± 100	24 ± 14	48 ± 28		$\textbf{0.09} \pm \textbf{0.07}$	24 ± 18		
Pannychia moseleyi	Dec. & Jan. mean	200 ± 73	27 ± 10	25 ± 10	82.9 ± 17.2	0.08 ± 0.07	23 ± 19	0.1 ± 0.0	27 ± 22
Chiridota sp.		150 ± 100	21 ± 14	12 ± 2.0	21.9 ± 16.3	0.07 ± 0.07	20 ± 19	0.3 ± 0.1	93 ± 50
Bathybembix bairdii		26 ± 18	4 ± 3	2.2 ± 0.2	25.7 ± 18.6	0.01 ± 0.01	4 ± 4	0.1 ± 0.0	14 ± 11
Т		380 ± 200	52 ± 27	39 ± 13		$\textbf{0.17} \pm \textbf{0.15}$	47 ± 41	0.5 ± 0.3	135 ± 83

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seafloor 234 Th_{xs} flux, based on seafloor inventories (Table 3). *Pannychia moseleyi* ingested the largest fraction of this flux, (25–27%), followed by *Chiridota* sp. (12–21%, Table 3).

Based on the *end-member mixing model*, *P. moseleyi*, *Chiridota* sp., and *B. bairdii* together ingested an average of $47 \pm 41\%$ of the average "fresh" particle flux over the Dec–Jan sampling times (Table 3). Again, *P. moseleyi* ingested the largest proportion. The percentage of zero-age sediment (i.e., sediment-trap material) in the gut, based on this model, was highest in *Pannychia* (83%), and was substantially lower (22–26%) for *Chiridota* sp. and *B. bairdii* (Table 3).

The estimated flux of sediment mass through all three animal populations in SCB was 0.5 g m⁻² d⁻¹ (Table 3). Assuming a dry bulk sediment density of \sim 0.2 g ml⁻¹ (Fornes *et al.* in prep.), these animals ingest a mass equivalent to the top 0.5 cm of SCB sediment in about 5 years.

5. Discussion

The gut contents of surface-deposit feeders in SCB were on average 24–176 times richer in 234 Th_{xs} than surface sediments (Fig. 2), suggesting that these animals were ingesting particles that had settled out of the water column within the last ~100 d. These results are in strong agreement with the age-dependent mixing hypothesis, which predicts that deep-sea deposit feeders feed preferentially on young, food-rich particles (C. Smith *et al.*, 1993).

Scotoplanes globosa and *P. moseleyi*, the elasipodid holothurians, had the highest weight-specific gut ²³⁴Th_{xs} activities in SCB, indicating that they ingest very fresh and presumably food-rich sediments. The high gut ²³⁴Th_{xs} activities no doubt result in part from feeding on a very thin veneer of surface sediment (i.e., $\ll 0.5$ cm thick) with mucous-covered tentacles. However, because elasipodids are quite mobile (moving at rates of 0.3 to 3 m hr⁻¹; Lafond, 1967; A. Smith *et al.*, 1997), this selective feeding likely also results in part from rapid exploitation of horizontal patchiness in recently deposited, food-rich (and ²³⁴Th_{xs}-rich) particles (cf. Yager *et al.*, 1993). In fact, *S. globosa*, and other species of elasipodids, do aggregate on patches of algal detritus and in shallow depressions where particle deposition may be enhanced (e.g., Smith and Hamilton, 1983); in addition, *S. globosa* is at least occasionally attracted to baitfalls (Pawson, 1976; C. Smith, 1985). These observations, in combination with our ²³⁴Th_{xs} and chl-*a* measurements, suggest that some elasipodids exploit meter-scale horizontal patchiness in "young," food-rich particles on time scales of hours to days.

While the gut ²³⁴Th_{xs} activities of *Chiridota* sp. and *B. bairdii* suggest substantial selectivity for relatively young particles (i.e., with apparent ages of 28–98 days), these species appeared to be much less selective than the elasipodid holothurians in SCB. For *Chiridota* sp., this lower selectivity may result from reduced ability to exploit horizontal patchiness due to much lower mobility. *Chiridota* sp. in SCB has a maximum measured burrowing speed of 10.0 cm d⁻¹ (mean = 6.5 cm d⁻¹) and while feeding extends a maximum of ~5 cm from its burrow (C. Smith 1992; G. Smith and C. R. Smith, unpublished data); thus, within a 24 hr period, a *Chiridota* individual can exploit patches no more

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than 15 cm apart. In contrast, the gastropod *B. bairdii* can move at speeds comparable to elasipodids (up to 0.9 m hr^{-1} ; Wheatcroft *et al.*, 1990) so its lower feeding selectivity is more likely to reflect differences in mechanisms of particle pickup and/or feeding strategies between snails and elasipodids (Lopez and Levinton, 1987; Penry and Jumars, 1987).

The 500-fold gut enrichment of chl-*a* in *P. moseleyi*, relative to surface sediments, reinforces the notion that freshly deposited phytodetritus may be an important food source for deep-sea animals (e.g., Billett *et al.*, 1988). Interestingly, the concentration of chl-*a* generally decreased along the gut of *P. moseleyi* (Fig. 5), suggesting that chl-*a* is degraded during gut passage. This contrasts with the result of Moore and Roberts (1994), who found little evidence of chl-*a* breakdown along the guts of three species of abyssal holothurians and suggested on this basis that phytoplankton material was not a direct source of nutrition to their holothurians.

On the Hawaiian slope, the surface-deposit feeding holothurians M. carnosa and P. retifer exhibited no enrichment over surface sediment (from a single core) in gut ²³⁴Th_{xs} activities. However, the 2.7-12 fold enrichment of chl-a in the guts of M. carnosa and *Phrissocystis* sp. suggests that these animals are indeed selective deposit feeders. What then explains the low gut enrichments of 234Th_{xs} in surface-deposit feeders on the Hawaiian slope relative to SCB? The answer most likely lies in differences in the depositional regimes between the two sites. SCB is a relatively quiescent basin, and our sampling site is located 14 km from the nearest slope. In contrast, the Hawaiian-slope site lies at the foot of a $\sim 22^{\circ}$ slope more than 1000-m in height (Fig. 1). The Hawaiian site is thus likely to experience substantial downslope transport of resuspended sediments (Edsall, 1975; Garcia, 1996). Such transport will entrain older bottom sediment into the water column where it can scavenge 234 Th and effectively decouple 234 Th_{xs} activity from time of particle settling from surface waters. Stated differently, the average ²³⁴Th_{xs}-rich particle on the Hawaiian slope is not likely to be a food-rich, recent arrival from the euphotic zone, and ²³⁴Th_{xs} activity thus would not be a good tracer of labile organic material. Enrichment of chl-a in the guts of the surface-deposit-feeders on the Hawaiian slope does suggest, however, that the deposit feeders feed selectively on phytodetritus newly deposited from the euphotic zone. Age-dependent mixing is, therefore, likely to occur on the Hawaiian slope, but we predict that ²³⁴Th_{xs} activity profiles will not reflect it.

Lauerman *et al.* (1997) found high 234 Th_{xs} activities in the guts of two species of holothurians (*Abyssocucumis abyssorum* and *Oneirophanta mutabilis*) from 4100-m depths off central California, with the gut sediments of these holothurians exhibiting mean apparent ages (relative to sediment-trap material) of 12–13 days. Thus, as predicted by the age-dependent mixing hypothesis (C. Smith *et al.*, 1993), the selective ingestion of young, 234 Th_{xs}-rich particles may be widespread among deposit feeders in quiescent deep-sea settings.

As expected, the guts of subsurface-deposit feeders were less enriched in 234 Th_{xs} than those of surface deposit feeders in SCB; *Molpadia* sp. had a mean gut 234 Th_{xs} activity at least 20-fold lower than any surface-deposit feeder (Fig. 2). The apparent age of sediments

in *Molpadia* sp. guts relative to surface sediment was ~50 days. Assuming that these holothurians are head-down conveyor-belt feeders like *Molpadia oolitica* (Rhoads and Young, 1971), they are subducting surface sediments to their feeding depth (~7 cm) within ~50 d, or they are feeding very selectively on 234 Th_{xs}-rich subsurface sediment. It seems unlikely that the molpadiids are intermittently feeding directly on surface sediments, because, to our knowledge, neither surface feeding nor down-burrow hoeing (as in Levin *et al.*, 1997) has ever been observed in molpadiids. Most likely, the gut 234 Th_{xs} activities result from some combination of rapid conveyor-belt feeding and particle selection at depth, because at least some molpadiids appear to be selective feeders (Self *et al.*, in prep).

With the *thorium-flux model*, we estimated that the combined populations of *P. moseleyi*, Chiridota sp., and B. bairdii in SCB ingested an average of 52% ($\pm 27\%$) of the daily sediment-trap flux of 234 Th_{xs}, and 39% (±13%) of the seafloor flux of 234 Th_{xs} based on sediment inventories (Table 3). The end-member mixing model yielded a very similar number, with 47% ($\pm 41\%$) of the daily flux of "fresh" particles into the sediment traps being consumed by these three species (Table 3). All three estimates have substantial uncertainties (Table 3). The major sources of error were (1) temporal variability in gut ²³⁴Th_{xs} activities (December means were 2–7 fold greater than those in January (Table 1)), and (2) spatial variability in the population densities (Table 2). We have no reason to think that either of these error sources is biased. One potential source of bias to the fluxconsumption estimates is the possibility of change in abundance of *Chiridota* sp. between 1987 (when time-lapse photographs were taken) and 1996. If *Chiridota* sp. were extremely rare in 1996, our estimates of the total flux consumed by the three megafaunal species would require revision downward by a factor of ~ 0.6 (i.e., roughly 30% of the daily sediment-trap flux of 234 Th_{xs} would still be consumed by SCB megafauna). However, the large numbers (thousands) of Chiridota sp. collected in each of our trawl samples (personal observations) suggests that these holothurians were still very abundant in 1995–1996. Based on the range in our two thorium-flux estimates (\pm one standard error), we conclude that P. moseleyi, Chiridota sp., and B. bairdii together consumed 39-52% of the daily 234 Th_{xs} flux to the SCB floor.

We assume that the flux of 234 Th_{xs} activity is a good proxy for labile POC reaching the SCB seafloor. Rapidly sinking, biogenic particles appear to carry most of the labile POC flux to the deep sea (e.g., Thiel *et al.*, 1989) and these particles also appear to be the main carriers of scavenged radionuclides such as 234 Th (Fowler and Knauer, 1986; C. Smith *et al.*, 1993). In quiescent deep-sea habitats where suitable measurements have been made, the fluxes of 234 Th_{xs} and POC are, as expected, very strongly correlated (e.g. Pope *et al.*, 1996; Lauerman *et al.*, 1997). Thus, ingestion of 39–52% of the daily 234 Th_{xs} flux implies that a similar percentage of the daily flux of POC is processed by *P. moseleyi, Chiridota* sp., and *B. bairdii* at the SCB floor.

Is it reasonable that 39-52% of the vertical flux of 234 Th_{xs} and POC to the SCB floor enters the guts of surface-deposit feeding megabenthos? Deep-sea holothurians (Khripounoff and Sibuet, 1980), and many shallow-water surface-deposit feeders (Lopez and

Levinton, 1987) exhibit assimilation efficiencies for bulk sedimentary POC on the order of 15%; this implies that *P. moseleyi, Chiridota* sp., and *B. bairdii* in aggregate may be assimilating and respiring roughly 6–8% of the vertical POC flux. Energetics studies of bathyal and abyssal benthic communities suggest that megafaunal deposit feeders, in aggregate, may easily respire this large a percentage of the seafloor POC flux (e.g., K. Smith, 1992). The injection rate of 234 Th_{xs} activity into the sediment column implied for *Chiridota* sp. also appears to be reasonable. Assuming that *Chiridota* sp. ingests 12% of the 234 Th_{xs} activity flux (Table 3) and defecates this activity into its burrow at a depth of ~1 cm (C. Smith, 1992), at steady state, at least 12% of the seafloor 234 Th_{xs} profile for SCB (Fornes, 1999) suggests that at least 12% of the seafloor inventory was indeed buried deeper than 0.5 cm. Thus, limited available data suggest that our calculated fluxes of POC and 234 Th_{xs} activity through megafaunal deposit feeders in SCB are reasonable.

Lauerman *et al.* (1997) used an end-member mixing model, based on ²¹⁰Pb in guts, surface sediments and sediment traps, to estimate that one holothurian species (*Abyssocucumis abyssorum*) processed 0.2–4% of the vertical particle flux to the abyssal seafloor off central California. Their estimate is substantially smaller than ours for at least two reasons. (1) *A. abyssorum* constituted only about 6% of the surface-deposit-feeding megafauna (holothurians and echinoids) at their site. (2) They used a gut throughput (or "residence" time) of 6 days for their holothurian, which is probably unreasonably long (Penry and Jumars, 1990; Billett, 1991). If a more reasonable gut throughput time of 24 hr is used, *A. abyssorum* is estimated to process 1.2–24% of the vertical particle flux, which is much more in line with our numbers for a holothurian with similar abundances, *P. moseleyi* (Table 3).

It should be noted that we base our particle consumption estimates in SCB on the flux of ²³⁴Th_{xs} activity through deposit-feeder guts (the *thorium-flux model*) rather than on an end-member mixing model using ²¹⁰Pb (half-life = 22 y), as in Lauerman et al. (1997). ²³⁴Th_{xs} is a more appropriate tracer for labile organic matter at the seafloor than ²¹⁰Pb because it decays on 100-day time scales similar to those of labile POC (e.g., Stephens et al., 1997). In contrast, ²¹⁰Pb integrates processes occurring over time scales of roughly 100 yr. The thorium-flux model also is a better approach than the end-member mixing *model* for at least two additional reasons. (1) Its primary assumption is that 234 Th_{xs} is a good seafloor tracer for labile POC, while the end-member mixing model has a number of additional untested assumptions (in particular, that animals feed on either "zero-age" or bulk sediments, see Section d. Flux consumption estimates). Thus, the more direct thorium-flux model is likely to yield more accurate estimates of the ingestion rates of labile POC. (2) In quiescent habitats, gut fluxes based on the thorium-flux model can be compared to seafloor fluxes derived from sediment inventories of ²³⁴Th_{xs}. This allows one to estimate the percentage of vertical flux consumed without deploying sediment traps, making the approach logistically simpler and independent of sediment-trap biases.

The finding that megafaunal deposit feeders may ingest a large fraction of the vertical

POC flux in SCB has important geochemical and ecological implications. The organic material ingested by deposit feeders is chemically altered (Plante et al., 1990), with some material assimilated for respiration and growth. Deep-sea holothurians, for example, have been shown to assimilate an average of 15% of the POC, 22% of the organic nitrogen, and 52% of the hydrolyzable amino acids from ingested deep-sea sediments (Khripounoff and Sibuet, 1980; Sibuet, 1988). The remainder is passed as feces, which are often still highly enriched in organic matter relative to surface sediment (e.g., Plante et al., 1990). Deposit-feeder feces appear to foster enhanced rates of microbial decomposition and solubalization of organic substances (e.g. Deming and Colwell, 1982; Plante et al., 1990), so the fecal deposits of surface-feeding megabenthos are likely to constitute major geochemical "hot spots" (cf. Aller, 1982). In addition to digestive alterations, *Chiridota* sp. appears to be an important bioturbator, subducting $\sim 12\%$ of the daily vertical POC flux to sediment depths of ~ 1 cm. The buried chiridotid fecal material is likely to be sequestered from the high microbial activity characteristic of the sediment-water interface (e.g., Webb and Montagna, 1993), promoting organic-matter preservation. Such subduction clearly may be regarded as age-dependent mixing, and may help to explain the high ²³⁴Th_{xs} mixing rates measured in Santa Catalina Basin (C. Smith et al., 1993).

The apparent control by megafauna of a large proportion of the food supply in SCB and off central California has profound implications for the entire benthic community. Subsurface-deposit-feeders and the relatively immobile sediment assemblage (bacteria through macrofauna) may have little access to roughly 50% of freshly deposited particles, with this material becoming available only after ingestion (and geochemical alteration) by the mobile megabenthos.

Fecal casts produced by mobile megafauna are likely to increase the patchiness of food resources, potentially promoting high species diversity in deep-sea sediments (Snelgrove *et al.*, 1992). If our results from SCB and those from central California (Lauerman *et al.*, 1997) are typical of the deep sea, then the mobile, surface-deposit-feeding megabenthos must substantially influence the structure, geochemistry and energetics of deep-sea ecosystems, and should be considered to be "keystone resource modifiers" (Levin *et al.*, 1997). The role of large mobile deposit feeders in modulating sediment geochemistry and energy flow clearly merits study in additional deep-sea habitats.

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REFERENCES

Aller, R. C. 1982. The effects of macrobenthos on chemical properties of marine sediments and overlying water, *in* Animal-Sediment Relations, P. L. McCall and M. J. S. Tevesz, eds., Plenum, NY, 53–101.

- 1990. Bioturbation and manganese cycling in hemipelagic sediments, *in* The Deep Sea Bed: Its Physics, Chemistry, and Biology, H. Charnock *et al.*, eds., Cambridge University Press, Cambridge, 51–68.
- Aller, R. C. and D. J. DeMaster. 1984. Estimates of particle flux and reworking at the deep-sea floor using Th-234/U-238 disequilibrium. Earth Planet Sci. Letts., 67, 308–318.
- Archer, D., S. Emerson and C. R. Smith. 1989. Direct measurement of the diffusive sublayer at the deep-sea floor using oxygen microelectrodes. Nature, 340, 623–626.
- Billet, D. S. M. 1991. Deep-sea holothurians. Oceanogr. Mar. Biol. Ann. Rev., 29, 259-317.
- Billet, D. S. M., C. Llewellyn and J. Watson. 1988. Are deep-sea holothurians selective feeders? *in*, Echinoderm Biology, R. Burke, P. Mladenov, P. Lambert, and R. Parsley, eds., Balkema, Rotterdam, 421–429.
- Chave, E. H. and A. T. Jones. 1991. Deep-water megafauna of the Kohala and Haleakala slopes, Alenuihaha Channel, Hawaii. Deep-Sea Res., *38*, 781–803.
- Deming, J. W. and R. R. Colwell. 1982. Barophilic bacteria associated with the digestive tract of abyssal holothurians. Appl. Environ. Microbiol., 44, 1222–1230.
- Dunbar, R. B., A. R. Leventer and W. L. Stockton. 1989. Biogenic sedimentation in McMurdo Sound, Antarctica. Mar. Geol., 85, 155–179.
- Edsall, D. J. 1975. Submarine Geology of Volcanic Ash Deposits: Age and Magmatic Composition of Hawaiian and Aleutian Tephra; Eocene to Recent. PhD dissertation, Columbia University, New York, NY.
- Emery, K. O. 1960. The Sea Off Southern California. John Wiley and Sons Inc., NY.
- Fornes, W. L. 1999. A radiochemical method for examining short-term sediment mixing processes in continental slope and deep sea sediments. Ph.D. dissertation, North Carolina State University, Raleigh, NC.
- Fowler, S. W. and G. A. Knauer. 1986. Role of large particles in the transport of elements and organic compounds through the oceanic water column. Prog. Oceanogr., *11*, 147–194.
- Gage, J. D. and P. A. Tyler. 1991. Deep-Sea Biology: A Natural History of Organisms at the Deep-Sea Floor. Cambridge University Press, 504 pp.
- Garcia, M. O. 1996. Turbidites from slope failure on Hawaiian volcanoes, *in* Volcano Instability on the Earth and Other Planets, McGuire, Jones, and Neuberg, eds., Geological Society Special Publication No. 110, 281–294.
- Hodges, J. L. and E. L. Lehmann. 1970. Basic Concepts of Probability and Statistics. Holden-Day Inc.
- Jeffrey, S. W. and G. M. Hallegraeff. 1987. Chlorophyllase distribution in ten classes of phytoplankton: a problem for chlorophyll analysis. Mar. Ecol. Progr. Ser., *35*, 293–304.
- Jumars, P. A. and R. A. Wheatcroft. 1989. Responses of benthos to changing food quality and quantity, with a focus on deposit feeding and bioturbation, *in* Productivity of the Ocean: Present and Past, W. H. Berger, V. S. Smetacek and G. Wefer, eds., John Wiley & Sons, 235–254.
- Karl, D. M. and R. Lukas. 1996. The Hawaii Ocean Time-series (HOT) program: background, rationale, and field implementation. Deep-Sea Res. II, *43*, 129–156.
- Khripounoff, A. and M. Sibuet. 1980. La nutrition d'echinodermes abyssaux: I. Alimentation des holothuries. Mar. Biol., 60, 17–26.
- LaFond, E. C. 1967. Movements of benthonic organisms and bottom currents as measured from the bathyscaphe *Trieste, in* Deep Sea Photography, J. B. Hershey, ed., Johns Hopkins Press, 295–302.
- Lauerman, L. M. L., J. M. Smoak, T. J. Shaw, W. S. Moore, and K. L. Smith Jr. 1997.²³⁴Th and ²¹⁰Pb evidence for rapid ingestion of settling particles by mobile epibenthic megafauna in the abyssal NE Pacific. Limnol. Oceanogr., *42*, 589–595.

- Levin, L., N. Blair, D. DeMaster, G. Plaia, W. Fornes, C. Martin and C. Thomas. 1997. Rapid subduction of organic matter by maldanid polychaetes on the North Carolina Slope. J. Mar. Res., 55, 595–611.
- Lopez, G. R. and J. S. Levinton. 1987. Ecology of deposit-feeding animals in marine sediment. Quart. Rev. Biol., 62, 235–260.
- McNeil, G. T. 1954. Photographic Methods: Problems and Solutions. Pitman, NY, 244 pp.
- Moore, H. M. and D. Roberts. 1994. Feeding strategies in abyssal holothurians, *in* Echinoderms Through Time, David, Guille, Feral and Roux, eds., Balkema, Rotterdam.
- Pawson, D. L. 1976. Some aspects of the biology of deep-sea echinoderms. Thalass. Jugoslav., *12*, 287–293.
- Penry, D. L. and P. A. Jumars. 1987. Modeling animal guts as chemical reactors. Am. Nat., 129, 69–96.
- 1990. Gut architecture, digestive constraints and feeding ecology of deposit-feeding and carnivorous polychaetes. Oecologia, *82*, 1–11.
- Plante, C. J. and P. A. Jumars. 1992. The microbial environment of marine deposit-feeder guts characterized via microelectrodes. Microbial Ecol., 23, 257–277.
- Plante, C. J., P. A. Jumars and J. A. Baross. 1990. Rapid bacterial growth in the hindgut of a marine deposit feeder. Microbiol Ecol., 18, 29–44.
- Pope, R. H. 1992. Particle mixing at two continental margin sites: a multitracer radiochemical approach. M.S. thesis, North Carolina State University, Raleigh, NC.
- Pope, R. H., D. J. DeMaster, C. R. Smith and H. Seltmann, Jr. 1996. Rapid bioturbation in Equatorial Pacific sediments: Evidence from excess Th-234 measurements. Deep-Sea Res. II, 43, 1339–1364.
- Rhoads, D. C. and D. K. Young. 1971. Animal-sediment relations in Cape Cod Bay, Massachusetts II. Reworking by *Molpadia oolitica* (Holothuroidea). Mar. Biol., *11*, 255–261.
- Roberts, D., H. Moore, B. Manship, G. Wolff, V. Santos, I. Horsfall, J. Patching and D. Eardly. 1996. Feeding strategies and impact of holothurians in the deep sea, *in* Irish Marine Science, B. F. Keenan and R. O'Connor, eds., Galway Univ. Press, 237–251.
- Sibuet, M. 1988. Structure des Peuplements Benthiques en Relation avec les Conditions Trophiques en Milieu Abussal dans l'Ocean Atlantique. Cas Particular des Echinoderms. These de Doctorat d'état es Sciences Naturelles, Universite des Pierre et Marie Curie, Paris.
- Smith, A., J. Matthiopoulos and I. G. Priede. 1997. Areal coverage of the ocean floor by the deep-sea elasipodid holothurian *Oneirophanta mutabilis:* estimates using systematic, random and directional search strategy simulations. Deep-Sea Res., 44, 477–486.
- Smith, C. R. 1985. Colonization studies in the deep sea: are results biased by experimental designs? in Proceedings of the Nineteenth European Marine Biology Symposium, P. E. Gibbs, ed., Cambridge University Press, 183–190.
- 1986. Nekton falls, low-intensity disturbance and community structure of infaunal benthos in the deep sea. J. Mar. Res., *44*, 567–600.
- 1992. Factors controlling bioturbation in deep-sea sediments and their relation to models of carbon diagenesis, *in* Deep-Sea Food Chains and the Global Carbon Cycle, G. T. Rowe and V. Pariente, eds., Kluwer, Dordrecht, Netherlands, 375–393.
- Smith, C. R. and S. C. Hamilton. 1983. Epibenthic megafauna of a bathyal basin off Southern California: patterns of abundance, biomass, and dispersion. Deep-Sea Res., 30, 907–928.
- Smith, C. R., D. J. Hoover, S. E. Doan, R. H. Pope, D. J. DeMaster, F. C. Dobbs, and M. A. Altabet. 1996. Phytodetritus at the abyssal seafloor across 10° of latitude in the central equatorial Pacific. Deep-Sea Res. II, 43, 1309–1338.
- Smith, C. R., P. A. Jumars and D. J. DeMaster. 1986. *In situ* studies of megafaunal mounds indicate rapid sediment turnover and community response at the deep-sea floor. Nature, *323*, 251–253.

- Smith, C. R., R. H. Pope, D. J. DeMaster and L. Magaard. 1993. Age-dependent mixing of deep-sea sediments. Geochim. Cosmochim. Acta, 57, 1473–1488.
- Smith, K. L., Jr. 1992. Benthic boundary layer communities and carbon cycling at abyssal depths in the central North Pacific. Limnol. Oceanogr., 37, 1034–1056.
- Smith, K. L., Jr., R. S. Kaufmann and W. W. Wakefield. 1993. Mobile megafaunal activity monitored with a time-lapse camera in the abyssal North Pacific. Deep-Sea Res., 40, 2307–2324.
- Snelgrove, P. V. R., J. F. Grassle and R. F. Petrecca. 1992. The role of food patches in maintaining high deep-sea species diversity: field experiments using hydrodynamically unbiased colonization trays. Limnol. Oceanogr., 37, 1543–1550.
- Stephens, M. P., D. C. Kadko, C. R. Smith and M. Latasa. 1997. Chlorophyll-*a* and phaeopigments as tracers of labile organic carbon at the central equatorial Pacific seafloor. Geochim. Cosmochim. Acta, *61*, 4605–4619.
- Strickland, J. D. H. and T. R. Parsons. 1972. A Practical Handbook of Seawater Analysis, Fisheries Research Board of Canada, 167 pp.
- Sun, M., R. C. Aller and C. Lee. 1991. Early diagenesis of chlorophyll *a* in Long Island Sound sediments: a measure of carbon flux and particle reworking. J. Mar. Res., *49*, 379–401.
- Taghon, G. L., A. R. M. Nowell and P. A. Jumars. 1984. Transport and breakdown of fecal pellets: Biological and sedimentological consequences. Limnol. Oceanogr., 29, 64–72.
- Thiel, H., O. Pfannkuche, G. Schriever, K. Lochte, A. J. Gooday, C. H. Hemleben, R. F. G. Mantoura, C. M. Turley, J. W. Patching and F. Riemann. 1989. Phytodetritus on the deep-sea floor in a central oceanic region of the Northeast Atlantic. Biol. Oceanogr., 6, 203–239.
- Thunnel, R. C., C. H. Pilskaln, E. Tappa and L. R. Sautter. 1994. Temporal variability in sediment fluxes in the San Pedro Basin, Southern California Bight. Cont. Shelf Res., *14*, 33–352.
- Webb, D. J. and P. A. Montagna. 1993. Initial burial and subsequent degradation of sedimented phytoplankton: relative impact of macro- and meiobenthos. J. Exp. Mar. Biol. Ecol., 166, 151–163.
- Wheatcroft, R. A. 1992. Experimental tests for particle size-dependant bioturbation in the deep ocean. Limnol. Oceanogr., 37, 90–104.
- Wheatcroft, R. A., P. A. Jumars, C. R. Smith and A. R. M. Nowell. 1990. A mechanistic view of the particulate biodiffusion coefficient: step lengths, rest periods and transport directions. J. Mar. Res., 48, 177–207.
- Yager, P. L., A. R. M. Nowell and P. A. Jumars. 1993. Enhanced deposition to pits: a local food source for benthos. J. Mar. Res., 48, 83–91.