

1 DIFFERENCES IN THE CHEMICAL COMPOSITION OF ORGANIC-WALLED
2 DINOFLAGELLATE RESTING CYSTS FROM PHOTOTROPHIC AND
3 HETEROTROPHIC DINOFLAGELLATES¹

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30 *Condensed running title*
31 Dinoflagellate cyst compositional differences

32 ABSTRACT

33 Dinoflagellates constitute a large proportion of the planktonic biomass from marine to
34 freshwater environments. Some species produce a preservable organic-walled resting cyst
35 (dinocyst) during the sexual phase of their life cycle that is an important link between the
36 organisms, the environment in which their parent motile theca grew, and the sedimentary
37 record. Despite their abundance and widespread usage as proxy indicators for environmental
38 conditions, there is a lack of knowledge regarding the dinocyst wall chemical composition. It
39 is likely that numerous factors, including phylogeny and life strategy, determine the cyst wall
40 chemistry. However, the extent to which this composition varies based on inherent
41 (phylogenetic) or variable (ecological) factors has not been studied.

42 To address this, we used micro-Fourier transform infrared (FTIR) spectroscopy to
43 analyze nine cyst species produced by either phototrophic or heterotrophic dinoflagellates
44 from the extant orders Gonyaulacales, Gymnodiniales and Peridiniales. Based on the
45 presence of characteristic functional groups, two significantly different cyst wall
46 compositions are observed that correspond to the dinoflagellate's nutritional strategy. The
47 dinocyst wall compositions analyzed appeared carbohydrate-based, but the cyst wall produced
48 by phototrophic dinoflagellates suggested a cellulose-like glucan, while heterotrophic forms
49 produced a nitrogen-rich glycan. This constitutes the first empirical evidence nutritional
50 strategy is related to different dinocyst wall chemistries. Our results indicated phylogeny was
51 less important for predicting composition than the nutritional strategy of the dinoflagellate,
52 suggesting potential for cyst wall chemistry to infer past nutritional strategies of extinct taxa
53 preserved in the sedimentary record.

54

55 *Keywords*

56 dinoflagellate cyst, dinosporin, heterotrophic, infrared spectroscopy, macromolecule,
57 nutritional strategy, phototrophic

58

59 *Abbreviations*

60 FTIR, Fourier transform infrared

61 DCM, dichloromethane

62 EtOH, ethanol

63 Pyr-GC-MS, pyrolysis gas chromatography-mass spectrometry

64

65 INTRODUCTION

66 Dinoflagellates are biflagellate, eukaryotic protists that comprise a large proportion of
67 planktonic biomass (Taylor 1987) and are therefore important components of marine and
68 freshwater ecosystems. Some species, as part of their sexual reproduction and in preparation
69 for a dormant period in their life cycle, produce an organic-walled, refractory resting cyst
70 (Wall and Dale 1968, Pfister and Anderson 1987, Head 1996) capable of being preserved in
71 the sediment record. The existence of two life cycle stages has resulted in the creation of two
72 taxonomic systems: one used by biologists and based on the morphology of living motile
73 cells, and the other developed by paleontologists and based on resting cyst morphology.
74 Dinoflagellate resting cysts (dinocysts) are important sources of environmental information
75 for modern marine (e.g., Rochon et al. 1999, Dale et al. 2002, Matthiessen et al. 2005,
76 Holzwarth et al. 2007, Pospelova et al. 2008, Bouimetarhan et al. 2009, Holzwarth et al. 2010,
77 Zonneveld et al. 2013, Bringué et al. 2013) and freshwater (e.g., Kouli et al. 2001, Tardio et
78 al. 2006, Leroy et al. 2009, McCarthy et al. 2011, Mertens et al. 2012) ecosystems, and are
79 also valuable for the reconstruction of oceanographic conditions (e.g., Reichart and Brinkhuis
80 2003, Pospelova et al. 2006, González et al. 2008). They therefore represent the link between
81 organisms, the environment, and the sedimentary record. However, dinocyst taxa react in
82 different ways to oxidation; cysts from the order Peridiniales show high sensitivity, while
83 cysts from the order Gonyaulacales demonstrate greater resistance, both in laboratory and
84 natural settings (Dale 1976, Zonneveld et al. 1997, 2001, Combourieu-Nebout et al. 1998,
85 McCarthy et al. 2000, Hopkins and McCarthy 2002, Mertens et al. 2009a). This suggests that
86 intrinsic properties of the dinocyst walls vary between the different orders, some physical,
87 such as differences in wall thickness or structure, while others are chemical. In this study, we
88 focused on potential differences in cyst wall chemistry.

89 To date, little is known regarding the macromolecular composition of dinocyst walls,
90 which is comprised of a refractory biomacromolecule referred to as dinosporin (Fensome et

91 al. 1993). Different compositions have been suggested, including mainly aromatic (Kokinos
92 et al. 1998), a mixture of aromatic and aliphatic (Hemsley et al. 1994), and mainly aliphatic
93 and heavily cross-linked (Versteegh et al. 2007). More recently, a carbohydrate-based
94 dinosporin has been suggested on the basis of Fourier transform infrared (FTIR) analyses and
95 pyrolyses of *Lingulodinium machaerophorum* (Deflandre and Cookson) Wall from culture
96 and surface sediments (Versteegh et al. 2012). Further micro-FTIR spectroscopic analyses of
97 late Paleocene dinocysts in the genus *Apectodinium* (Costa and Downie) Lentin and Williams
98 agreed with this interpretation but also displayed considerable differences in composition
99 between species of the same genus (Bogus et al. 2012). Therefore, it seems that phylogeny
100 could be a factor contributing to dinosporin compositional differences. However, there are
101 also other factors with significant potential to influence dinosporin composition, including
102 differences in the compounds available within the motile dinoflagellate cell which builds the
103 cyst, and differences resulting from post-depositional alteration. It is assumed that
104 dinoflagellates build their cysts from material readily available within the cell (Kokinos and
105 Anderson 1995, Hallet, 1999), which was suggested by TEM studies of cyst formation in the
106 freshwater species *Ceratium hirundinella* (Müller) Dujardin (Chapman et al. 1982).
107 Furthermore, there is clear evidence that dinocyst morphology can vary with differing
108 environmental conditions (e.g., Hallet 1999, Ellegaard 2000, Lewis et al. 2001, Ellegaard et
109 al. 2002, Zonneveld and Susek 2007, Mertens et al. 2009b) as well as evidence that some
110 dinoflagellates, such as *Karlodinium veneficum* (Ballantine) Larsen and *Alexandrium*
111 *minutum* Halim (Fuentes-Grünewald et al. 2009, 2012) can adapt their cellular composition to
112 different environmental conditions. However, it is not known whether there is a link between
113 chemical compounds available within the motile cell and dinosporin composition. One way
114 to investigate potential impacts on dinosporin chemistry is to analyze dinocysts produced by
115 dinoflagellates that have very different life strategies, such as nutritional strategies
116 (phototrophy versus heterotrophy). Additionally, diagenetic processes, such as natural

117 vulcanization, have been shown to chemically alter the cyst walls of *Thalassiphora pelagica*
118 (Eisenack) Benedek and Gocht emend. (Versteegh et al. 2007) and raised the question as to
119 what extent the differences observed between *Apectodinium* species were influenced by
120 differential diagenesis (Bogus et al. 2012). However, we circumvented the diagenetic factor
121 by utilizing specimens isolated from modern marine and lacustrine surface sediments where
122 the exposure to diagenetic processes is short and thus expected to be small.

123 We investigated differences in the dinocyst wall chemistry in species from different
124 lineages (the orders Gonyaulacales, Gymnodiniales, and Peridinales). The main objectives
125 were to determine the primary factors that affect dinosporin composition and to address the
126 extent to which a particular dinocyst composition was monophyletic, i.e., different
127 dinoflagellate lineages produce different wall compositions, or polyphyletic and subject to
128 variations resulting from different nutritional strategies, since these are known to be
129 polyphyletic. The dinocysts were analyzed using micro-FTIR spectroscopy. This versatile
130 technique has already been successfully used to identify compounds in the complex
131 biopolymers of dinocyst wall layers of optically identified and individually isolated
132 specimens (e.g., Kokinos et al. 1998, Versteegh et al. 2007, 2012, Bogus et al. 2012).

133

134 MATERIAL AND METHODS

135 *Sample treatment*

136 Samples were retrieved from marine and lacustrine locations (Fig. 1, Table 1). Surface
137 sediment samples from the marine Benguela upwelling region (GeoB 2341, 4804), off the
138 coast of northwest Africa (GeoB 6010), and freshwater Honey Harbour in Lake Huron,
139 Ontario, Canada (SV5-C) were briefly ultrasonicated (< 60 s) in Milli-Q water to disaggregate
140 particles attached the cyst walls, sieved over a 50 µm nylon mesh sieve and retained on a 20
141 µm precision sieve (Storck Veco #317). Other specimens of empty dinocysts, previously
142 isolated for germination experiments, were analyzed from the Wadden Sea (NW Germany)

143 and Omura Bay (Kyushu, Japan). The various geographic locations, environmental settings
144 and isolation procedure (i.e. empty cysts isolated from the sediment matrix or empty cysts
145 utilized shortly after hatching) were used to verify that these different situations did not
146 overprint the cyst wall signal.

147 To remove extraneous apolar contaminants, all material was soaked in ethanol (EtOH)
148 or dichloromethane (DCM) for at least 30 min. To remove extraneous water soluble
149 molecules, all material was then rinsed three times with Milli-Q water. Similar to the
150 procedure described in Versteegh et al. (2007) and Bogus et al. (2012), individual specimens
151 with no visibly attached particles were identified to species level with a light microscope,
152 isolated with a micropipette, transferred to either a salt (NaCl) plate or an Au-coated mirror,
153 dried overnight at 60 °C, and analyzed immediately.

154 In total, specimens from nine extant dinocyst species from the orders Gonyaulacales,
155 Gymnodiniales, and Peridinales were isolated. The cyst species are produced by either
156 phototrophic or heterotrophic dinoflagellates. Cyst names are used in the descriptions. At
157 least three specimens of each species were picked and analyzed, and in the case of two
158 species, specimens from two different locations were analyzed (Table 2).

159

160 *Micro-Fourier transform infrared spectroscopy*

161 Infrared spectra of specimens from GeoB 2341, GeoB 4804, and GeoB 6010 were recorded
162 with a Nicolet FT-IR spectrometer coupled to a Nicplan microscope, a Protégé™ 460 optical
163 bench, a mercury cadmium telluride (MCT)- A detector cooled with liquid N₂, Ever-Glo
164 source, and a KBr beamsplitter. The adjustable apertures (upper and lower) were set at a
165 constant area of 15 x 15 μm. Two hundred and fifty-six scans at 8 cm⁻¹ resolution were
166 obtained in transmission mode over a spectral range of 4000-650 cm⁻¹. All other specimens
167 were analyzed with a BRUKER IFS 66v coupled to an IR Scope II equipped with a MCT
168 detector cooled with liquid N₂ and KBr beamsplitter. Two hundred and fifty-six scans at 4

169 cm^{-1} resolution were recorded in reflection mode over a spectral range of 4000-650 cm^{-1} . The
170 replicability of FTIR analysis using these two different devices was checked through the use
171 of a chitin standard (Sigma, Lot 59F7265). The spectra proved consistent and are displayed
172 after subtracting the background of air and the bare plates (NaCl [Nicolet]; Au [Bruker]), and
173 baseline correction. Assignments of the main IR absorptions to chemical bonds (Table 3)
174 were based on Colthup et al. (1990), Coates (2000), and additional published literature.

175

176 RESULTS AND DISCUSSION

177 *Evidence for carbohydrate-based dinosporins*

178 For all species, the region between 3600-3000 cm^{-1} showed a strong and broad absorption
179 with a maximum near 3350 cm^{-1} (OH stretch) and, sometimes, a shoulder near 3270 cm^{-1}
180 and/or 3100 cm^{-1} (Table 3, Figs. 2 and 3). Absorptions were relatively weak in the 3000-2775
181 cm^{-1} region (CH stretching), except for the spectrum of *Impagidinium patulum* (Wall) Stover
182 and Evitt, which exhibited two stronger absorptions at 2925 cm^{-1} and 2860 cm^{-1} . Since this
183 was observed in all *I. patulum* specimens, it is likely an intrinsic component. The relatively
184 weak absorptions in this region suggested a minor contribution from CH_2 and CH_3 groups.
185 There was a pattern of four absorptions between 1200-1030 cm^{-1} that are highly indicative of
186 C-O stretching and the deformation vibrations of sugar rings: 1160 cm^{-1} (C-O-C asymmetric
187 vibration), 1110 cm^{-1} (glucose ring stretch), 1060 cm^{-1} (C-O stretch), and 1030 cm^{-1} (C-O
188 stretch). This series was most apparent in *Dubridinium caperatum* Reid, *Spiniferites*
189 *pachydermus* (Rossignol) Reid, *Tuberculodinium vancampoe* (Rossignol) Wall, cysts of
190 *Polykrikos schwartzii* Bütschli, and cysts of *Peridinium wisconsinense* Eddy. The remaining
191 species, *Operculodinium centrocarpum* sensu Wall and Dale, *I. patulum*, *Brigantedinium* spp.
192 Reid, and cysts of *Polykrikos kofoidii* Chatton, exhibited most of these absorptions; however,
193 the appearance of this region either differed by containing less definition, as in *O.*

194 *centrocarpum*, *I. patulum*, and *Brigantedinium* spp., or by appearing shifted, as in cysts of *P.*
195 *kofoidii*. This variability is discussed further in subsequent sections.

196 Despite the variability exhibited, the absorptions between 1200-1030 cm^{-1} , together
197 with the diagnostic absorptions at 896-902 cm^{-1} , provide a strong argument that a
198 carbohydrate with a β -glycosidic linkage (Kačuráková and Wilson 2001, Cárdenas et al.
199 2004, Versteegh et al. 2012) forms the backbone of dinosporins. The cell walls of the motile
200 stage of the freshwater dinoflagellate *Peridinium westii* Lemmermann were shown to contain
201 both β -1 \rightarrow 3 and β -1 \rightarrow 4 glycosidic bonds (Nevo and Sharon 1969). Unfortunately, the FTIR
202 spectra of β -1 \rightarrow 3 linked glucans (e.g., Furuhashi et al. 2009) can be similar to those with β -
203 1 \rightarrow 4 linkages (e.g., Pandey 1999) that are also apparent in the highly resolved 1000-600 cm^{-1}
204 region described by Barker et al. (1954). It is therefore not currently possible to definitively
205 determine the type of β -glycosidic linkage in the dinosporin macromolecule. Regardless, the
206 analyzed species demonstrate carbohydrate-based dinosporins and complement the
207 interpretation of a carbohydrate backbone from FTIR spectroscopy and pyrolysis gas-
208 chromatography mass spectrometry of culture- and sediment-derived *Lingulodinium*
209 *machaerophorum* (Versteegh et al. 2012) as well as FTIR analyses of the extinct genus
210 *Apectodinium* (Bogus et al. 2012).

211

212 *Evidence for two broadly different dinosporin groups*

213 Regardless of the evidence for a carbohydrate-based dinosporin composition in all dinocyst
214 species, it appears that at least two broadly different types of dinosporin occur. This
215 distinction was based on the pattern of absorptions between 1850-650 cm^{-1} (Figs. 2 and 3) and
216 evidence of the main functional groups present (Table 3). The first dinosporin group, here
217 referred to as Group I, includes *Impagidinium patulum*, *Operculodinium centrocarpum*,
218 *Spiniferites pachydermus*, *Tuberculodinium vancampoe* (all Gonyaulacales) and the
219 freshwater cysts of *Peridinium wisconsinense* (Peridinales; Fig. 2). In this group,

220 absorptions between 1200-1030 cm^{-1} are dominant, absorptions between 1470-1200 cm^{-1}
221 (CH_2 and CH_3 bending and rocking; OH in-plane deformation vibrations) are second in
222 amplitude, and those between 1850-1600 cm^{-1} (adsorbed OH; conjugated C=O bonds) are the
223 weakest. The only exception was in the spectrum of *T. vancampoeae*, where the absorption at
224 1600 cm^{-1} is almost as strong as the 1200-1030 cm^{-1} region.

225 There was spectral variability within Group I, such as the changes in relative intensity
226 between the 1640 cm^{-1} and 1600 cm^{-1} absorptions (Fig. 2). The most prominent 1640 cm^{-1}
227 absorption is exhibited in the *S. pachydermus* spectrum, whereas it is either a shoulder or a
228 weak peak in the other species. This pattern is reversed for the absorption at 1600 cm^{-1} where
229 it is only a shoulder in *S. pachydermus*, but a stronger signal in *O. centrocarpum* and,
230 especially, *T. vancampoeae*. A strong peak at 1600 cm^{-1} could suggest more ester bonds in
231 those species' cyst walls (e.g., Yuen et al. 2009). However, minor absorptions between 3000-
232 2800 cm^{-1} and 1470-1350 cm^{-1} imply that methylene and methyl groups, and thus ester bonds,
233 cannot be responsible for a more pronounced absorption at 1600 cm^{-1} . It is also possible that
234 this represents the influence of aromatic skeletal vibrations, as shown in lignin (Pandey 1999),
235 but there is no clear evidence for aromatic vibrations ($\sim 1500 \text{ cm}^{-1}$) in any of the dinocyst
236 spectra. Therefore, at this point, a more definitive explanation of the 1600 cm^{-1} absorption is
237 not available.

238 There was also variability in the deformation pattern between 1200-850 cm^{-1} . This
239 region is comprised of four separate, defined absorptions in the cysts of *P. wisconsinense*, *S.*
240 *pachydermus*, and *T. vancampoeae* that closely match the spectrum of the β -linked glucan,
241 cellulose (Fig. 2) as well as the previously published spectrum for *Lingulodinium*
242 *machaerophorum* (Versteegh et al. 2012). In fact, the spectrum of *S. pachydermus* was so
243 remarkably similar to cellulose overall that each of the absorptions observed for *S.*
244 *pachydermus* can easily be assigned using the cellulose spectrum (Pandey 1999). The lack of
245 this absorption series in *I. patulum* and *O. centrocarpum*, particularly the peak at 1112 cm^{-1} ,

246 and the existence of a broader absorption centered at 1060 cm^{-1} may indicate that glucose is
247 not the only sugar monomer present. While cellulose is the best known β -glucan (Aspinall
248 1983) and is described as the primary material comprising the theca of dinoflagellates (Sekida
249 et al. 2004), other non-cellulosic β -glucans, such as mannan, are common and well-
250 documented components in plant (e.g., Kačuráková and Wilson 2001, Burton and Fincher
251 2009) and algal cell walls (e.g., Frei and Preston 1964, Stone 2009), including the motile
252 dinoflagellate *Peridinium westii* (Nevo and Sharon 1969). Thus, it is likely that non-
253 cellulosic β -glucans also contribute to the carbohydrate signal of dinosporins in this group. In
254 general, we propose that a spectral signal indicating β -glucans represents the signature
255 composition of these dinocyst species.

256 Dinocysts in the second group (Group II) consist of *Brigantedinium* spp., *Dubridinium*
257 *caperatum* (both Peridinales), and cysts of *Polykrikos schwartzii* and *P. kofoidii*
258 (Gymnodinales). These species exhibited a different spectral pattern of relative absorption
259 strength, demonstrated greater heterogeneity with respect to the relative intensity of each of
260 the regions, and, most significantly, included evidence for nitrogen (N)-containing functional
261 groups (Fig. 3, Table 3). The region between $1850\text{-}1500\text{ cm}^{-1}$ dominated in *D. caperatum* and
262 cysts of *P. kofoidii*, but it was less intense in cysts of *P. schwartzii* and *Brigantedinium* spp.
263 Within this region, maxima occurred especially between $1585\text{-}1560\text{ cm}^{-1}$, and there was a
264 clear shoulder at 1660 cm^{-1} . Absorptions between $1585\text{-}1560\text{ cm}^{-1}$ are characteristic of amide
265 II bonds (CN stretching and NH bending), while the shoulder at 1660 cm^{-1} in all of the
266 species probably reflects amide I bonds, which is the result of the influence of hydrogen
267 bonding ($\text{C}=\text{O}\cdots\text{H}-\text{N}$; Cárdenas et al. 2004). The area between $1500\text{-}1200\text{ cm}^{-1}$ dominated
268 only in *Brigantedinium* spp., although all of the Group II dinocysts showed a small absorption
269 near 1255 cm^{-1} (NH bending). The absorptions between $1420\text{-}1370\text{ cm}^{-1}$ reflect CH bending,
270 and the absorption at 1312 cm^{-1} indicates CN stretching and NH bending (amide III). Further
271 evidence for the presence of nitrogen included a shoulder at 3100 cm^{-1} (NH stretching),

272 further NH stretching absorptions ($\sim 3268\text{ cm}^{-1}$) likely encompassed within the broad OH
273 stretching region ($3600\text{-}3000\text{ cm}^{-1}$) and manifested by a shifting of the peak center (relative to
274 Group I), and a small peak at 698 cm^{-1} (NH wagging; amide V). Absorptions between 1200-
275 1030 cm^{-1} (C-O stretching) account for a much smaller proportion of the total absorptions
276 from $1850\text{-}830\text{ cm}^{-1}$ with the exception of cysts of *P. schwartzii*, where the absorptions in this
277 region dominated and clearly demonstrated the four separate absorptions characteristic of
278 sugar ring vibrations present in many of the Group I dinocysts. In combination, the spectral
279 evidence in Group II included many absorptions that are typically seen in the spectrum of the
280 polysaccharide chitin (Fig. 3a; e.g., Cárdenas et al. 2004) as well as peptides (e.g.,
281 Venyaminov and Kalnin 1990). The evidence of N-containing functional groups may reflect
282 single amino acids, as is the case for chitin, or more complex, proteinaceous material such as
283 (oligo)peptides. The combination of the carbohydrate evidence, together with the amide bond
284 evidence, suggest a dinosporin composition based on a more chitin-like glycan (Stankiewicz
285 et al. 1998, Kačuráková et al. 1999, Cárdenas et al. 2004) or even a chitin-glucan complex
286 (Šandula et al. 1999). The exhibited in-group variability could indicate different sugar
287 moieties, as postulated for Group I, suggest contributions of different amino acids
288 (Venyaminov and Kalnin 1990) or chitins (i.e. α or β , Cárdenas et al. 2004), and/or reflect
289 varying ratios of N-containing functional groups to the carbohydrate backbone.

290

291 *Explanation of compositional differences*

292 Our results demonstrate considerable spectral and, thus, compositional diversity among the
293 dinocyst species that indicate dinosporin is a chemically heterogeneous compound. The most
294 fundamental distinction between the dinosporins is the inclusion of N-containing functional
295 groups in Group II and we now explore the reasons for this difference in composition. Each
296 of the dinocyst species in Group I is produced by phototrophic dinoflagellates, while the
297 Group II dinocysts are produced by heterotrophic dinoflagellates (Table 2). The heterotrophic

298 species studied here prey upon a variety of dinoflagellates and diatoms (e.g., Jacobson and
299 Anderson 1986, Matsuoka et al. 2000, Naustvoll 2000, Menden-Deuer et al. 2005).
300 Therefore, the origin of the amide groups in Group II dinosporins may be from predation by
301 the dinoflagellates, which leads to an accumulation of N-rich compounds (i.e., proteinaceous
302 compounds) within the cell as a result of prey digestion. Many heterotrophic organisms,
303 whose growth is energy limited, produce cell coverings that contain both amino acids and
304 sugars like peptidoglycans (bacteria) and chitin (arthropods and fungi) because both types of
305 compounds are abundant in prey and therefore energetically favorable to use. On the other
306 hand, phototrophic organisms are not energy limited, but nutrient limited (i.e. nitrogen and
307 phosphorus). Therefore, it is not energetically favorable for them to utilize these limited
308 nutrients to build a metabolically inactive cell covering, but rather to incorporate the products
309 of photosynthesis (e.g., Thornton et al. 1999, Wotton 2004, Ellegaard et al. 2013). As
310 dinocyst walls are assumed to be constructed using compounds from within the dinoflagellate
311 cell (Kokinos and Anderson 1995), the incorporation of photosynthetic products into a Group
312 I dinocyst would account for a glucan backbone, while the predominance of proteinaceous
313 material would lead to the inclusion of N-containing functional groups into Group II
314 dinocysts. In both cases, dinoflagellates thus use compounds that are in excess and
315 energetically favorable, determined by their nutritional strategy.

316 Compositional differences between phototrophic and heterotrophic dinoflagellates and
317 their cysts have previously been suggested by studies investigating the autofluorescence of
318 the two groups. Fluorescence microscopy was used to distinguish motile photosynthetic
319 dinoflagellates from heterotrophic ones (Lessard and Swift 1986). Additional work with
320 fossil and recent heterotrophic dinocysts demonstrated they do not exhibit autofluorescence
321 (Brenner and Biebow 2001) and this absence has been used to infer a heterotrophic ecology in
322 newly described dinocyst species (Verleye et al. 2011). Our results support these findings by

323 providing the first description of differences in chemical composition between phototrophic
324 and heterotrophic dinocysts.

325 In contrast to the expectation that dinocyst composition might solely exhibit
326 characteristic phylogenetic differences, a clear division along phylogenetic lines is absent.
327 Most of the Group I dinocysts are in the order Gonyaulacales, while Group II dinocysts are
328 either Peridinales or Gymnodinales; however, cysts of *Peridinium wisconsinense* exhibit
329 spectra firmly positioned in Group I despite being in the order Peridinales. Both
330 compositional groups are polyphyletic when compared to dinoflagellate phylogeny (Fensome
331 et al. 1993, Taylor 2004, Zhang et al. 2007, Hoppenrath and Leander 2010, Orr et al. 2012;
332 Fig. 4). Therefore, we propose that nutritional strategy rather than phylogeny is the primary
333 factor identified which determines cyst wall composition. However, related species tend to
334 have related life strategies so some covariance with species phylogeny is expected. This is
335 most evident in Group II where the spectra of the *Polykrikos* species are more similar in terms
336 of relative absorption intensities than to the other taxa investigated in this group (Fig. 3).

337 We cannot exclude the possibility that some of the differences observed result from
338 taphonomic and environmental heterogeneity and/or preparative processes. Taphonomic
339 processes, such as sulfurization in anoxic environments (Kok et al. 2000, van Dongen et al.
340 2003, Versteegh et al. 2007) or oxidative polymerization in oxygenated settings (Versteegh et
341 al. 2004, Gupta et al. 2006, de Leeuw 2007), can overprint the original biomacromolecular
342 signal. Even though the analyzed specimens were isolated from surface sediments, they may
343 have undergone some early diagenetic alteration that led to the loss of the most easily
344 degradable components. Diagenetic effects on dinosporin composition are not well known,
345 but Versteegh et al. (2007) showed an initial preferential loss of oxygen-containing functional
346 groups as well as an increase in aliphatic CH_x bonds during diagenesis. All of our measured
347 specimens demonstrated strong absorptions for oxygen-containing functional groups and

348 relatively weaker absorptions for aliphatic CH_x, suggesting that diagenesis as a whole has not
349 significantly affected the cyst wall chemistry.

350 Some of the variability seen between and within the two groups may also be explained
351 by the presence, absence and/or thickness of layers with different contributions of
352 carbohydrates and amino acids, or it may be that different proportions of these compounds
353 remain attached to the cyst wall after hatching or migrate onto the cyst wall from the
354 sediment. However, there was clear consistency between spectra of the same dinocyst species
355 measured from more than one location, *Operculodinium centrocarpum* and cysts of
356 *Polykrikos schwartzii* (Table 2, Figs. 2, 3 [gray lines]). The distinction between Group I and
357 II dinosporins was also consistent between empty cysts of *Tuberculodinium vancampoae* and
358 *Dubridinium caperatum* measured shortly after hatching from. The maintenance of this
359 distinction (i.e. the absence of nitrogen-containing functional group evidence in *T.*
360 *vancampoae*) indicates little effect, if any, from leftover cell material after sample treatment.

361 In terms of preparative alteration, the specimens analyzed all represent visibly and
362 chemically clean, but thermally untreated material. Therefore, the origin of the amide bonds
363 cannot be artifacts, such as melanoidin-like polymers, produced during preparation (Allard et
364 al. 1997, 1998). Cleaning involved gentle ultrasonic treatment (< 60 s; Mertens et al. 2009a)
365 and solvent extraction, specifically washing with Milli-Q water to remove water soluble
366 contaminants, and DCM/EtOH to remove apolar contaminants. The use of solvent extraction
367 alone, as opposed to its use in conjunction with acid and/or base hydrolysis (e.g., Kokinos et
368 al. 1998, Versteegh et al. 2012), was used for a couple of reasons. First, even a short treatment
369 with a base selectively destroys peridinialean cysts (e.g., Dale, 1976, Mertens et al. 2009a).
370 Additionally, a comparison between extracted and hydrolyzed cultured *Lingulodinium*
371 *machaerophorum* cysts and extracted sediment-derived cysts showed remarkably similar
372 spectra (Versteegh et al. 2012), suggesting that hydrolysis was not necessary to render the
373 cysts chemically clean. The *L. machaerophorum* spectra are very consistent with the Group I

374 spectra (Fig. 2). It may not be possible to fully discount effects of invisible material on/in cyst
375 walls insoluble in water and organic solvents on resulting spectra, particularly for species
376 with significant surface ornamentation such as *O. centrocarpum*, *T. vancampoae*, *Spiniferites*
377 *pachydermus* and the *Polykrikos* species, but the lack of evidence for an overprint indicates
378 this is unlikely; instead, it suggests that preparing cysts with solvent extraction has the
379 potential for wide applicability as even delicate cysts with lower resistance to harsh chemical
380 treatments will be able to be reliably analyzed.

381 The aforementioned observations indicating (1) little to no diagenetic effects, (2)
382 consistency between the same species analyzed from multiple locations as well as the
383 maintenance of the group-wise distinction between different species of freshly hatched cysts,
384 and (3) the similarity of the Group 1 spectra to spectra from hydrolyzed *L. machaerophorum*
385 cysts argue against any significant overprint from diagenesis, environment and sample
386 preparation.

387

388 *Implications for (paleo)environmental studies*

389 A determination of two broadly different dinosporin types was based on the FTIR spectral
390 patterns and functional groups present, and attributed to nutritional strategy. The spectral
391 variability, particularly with regards to the relative strength of the individual absorptions,
392 suggests that the story is likely more complicated given the large variety in nutritional
393 strategies amongst dinoflagellates (Schnepf and Elbrächter 1993) as well as previous evidence
394 that aspects of dinosporin composition are also likely taxon-specific (de Leeuw et al. 2006,
395 Bogus et al. 2012, Figs. 2, 3). So, the spectral variability within the two groups does not
396 exclude a phylogenetic component in dinosporin synthesis, only that it is obscured in this case
397 by the complex interplay between the dinoflagellate and its environment. Thus, it may be
398 more accurate to define dinosporin overall as a suite of chemically distinct but related
399 carbohydrate-based biomacromolecules. The most apparent distinction in a broad sense is the

400 inclusion of amide groups into some species' dinosporins, resulting from heterotrophy, and in
401 this sense dinosporin does not follow a strict delineation with phylogeny (Fig. 4). Therefore,
402 based on the evidence, we suggest FTIR spectra of dinocysts have the potential to assess
403 (paleo)nutritional strategies. This may prove especially interesting for species that may
404 change their nutritional strategy in response to environmental conditions.

405 Very important subsequent steps are to expand the number of investigated taxa in
406 general and through the sedimentary record to robustly determine the predictive value of
407 dinosporin composition for (paleo)ecology. Even though carbohydrates are the most abundant
408 form of biomass on earth (Kurita 2006), they are usually considered labile compounds (e.g.,
409 de Leeuw and Largeau 1993, Arnosti 1995) with preservation determined by the lability of
410 individual sugar monomers (Moers et al. 1994, Marchand et al. 2009) and diagenetic
411 alteration, such as macromolecular skeletal rearrangements (Almendros et al. 1997). There is
412 evidence for carbohydrate preservation relatively far back in the sedimentary record
413 ([Miocene] e.g., Lechien et al. 2006), including well-preserved glycolipids ([Eocene]
414 Bauersachs et al. 2010) as well as spectra demonstrating good preservation of Group I
415 compositional signals ([late Paleocene] Bogus et al. 2012). Dinospirin is likely to involve
416 some inherent significant structural differences to straightforward carbohydrates, such as
417 having a more highly cross-linked backbone (Versteegh et al. 2012), which explains the
418 generally high preservation potential of dinocysts. These structural differences are
419 unfortunately not resolvable in FTIR spectra. However, the *L. machaerophorum* cysts were
420 analyzed for structural information, which indicated a carbohydrate-based polymer
421 (Versteegh et al. 2012). Due to the strong similarities between *L. machaerophorum* and
422 Group I spectra (Fig. 2), it is probable that the cyst macromolecule is comparable.

423 Selective preservation of dinocysts in the sedimentary record has been suggested
424 (Zonneveld et al. 1997, 2001, Combourieu-Nebout et al. 1998, McCarthy et al. 2000,
425 Versteegh et al. 2010), although it is likely not a straightforward process (Reichart and

426 Brinkhuis 2003). Based on these studies peridinialean cysts have been considered more
427 sensitive to oxidation than gonyaulacalean cysts, which is supported by laboratory evidence
428 that peridinialean cysts are destroyed after harsh base treatment (e.g., Dale 1976, Hopkins and
429 McCarthy 2002, Mertens et al. 2009a). However, based on dinosporin composition, we
430 suggest that sensitivity to oxidation is more likely predicated on dinosporin composition. For
431 example, the peridinialean cysts of *P. wisconsinense* withstand treatment by harsh bases
432 during palynological preparation unlike many marine peridinialean species; there are also
433 indications they are more prevalent in sedimentary successions than previously recognized
434 (e.g., Miller et al. 1982, Zippi et al. 1990, McCarthy et al. 2011, Mertens et al. 2012).
435 Therefore, the Group I dinocysts may be inherently more resistant while the Group II
436 dinocysts are more heavily reliant on additional diagenetic factors, such as sulfurization or
437 skeletal rearrangement, to facilitate preservation. It is also interesting to note that the Group
438 II cysts are brown in color, while Group I dinocysts are transparent, suggesting the possibility
439 that the inclusion of nitrogen-containing functional groups and/or different pigments may also
440 contribute to increased oxidative sensitivity. Regardless, these compositional differences
441 between dinosporins provide the first empirical evidence that differences in preservation
442 potential may be related to dinocyst wall chemistry.

443

444 CONCLUSIONS

445 Micro-FTIR spectroscopic analyses of recent dinoflagellate resting cyst (dinocyst) walls from
446 three large orders (Gonyaulacales, Gymnodiniales, Peridinales) showed that all of the
447 dinocysts had a carbohydrate-based macromolecular composition. Spectral and
448 compositional variability indicated that the cyst wall macromolecule (dinosporin) is a
449 collection of chemically distinguishable but related biomacromolecules determined by the
450 interaction of the cyst-producing dinoflagellate with its environment. Two groups were
451 defined on the basis of this spectral variability. The first group encompasses dinocysts

452 produced by phototrophic dinoflagellates, which includes all the gonyaulacalean cysts and the
453 peridinialean cysts of *Peridinium wisconsinense*. Their dinosporin has a β -glucan backbone,
454 which concurs with a previous study of *Lingulodinium machaerophorum* (Versteegh et al.
455 2012). The second group consists of dinocysts produced by heterotrophic dinoflagellates,
456 including the remainder of the peridinialean and all of the gymnodinialean cysts studied.
457 They showed the presence of amide bonds within their dinosporins. These systematic
458 differences reflect the nutritional strategy rather than the phylogeny of the cyst-producing
459 dinoflagellate and may result from the incorporation of chemical compounds in excess within
460 the cell into dinosporins. Thus, the nitrogen-containing functional groups present in the
461 heterotrophic dinosporins may originate from the digestion of prey and the resulting
462 proteinaceous, nitrogen-rich material. The nutritional strategies of the cyst-producing
463 dinoflagellates were polyphyletic, as demonstrated by the grouping of the peridinialean cysts
464 of *P. wisconsinense* with the other phototrophic species (all Gonyaulacales) as well as the
465 grouping of remaining Peridinales and Gymnodinales cysts. As nutritional strategy, not
466 phylogeny, appears to be the primary factor determining dinosporin composition, cyst wall
467 chemistry may show potential as a paleoecological proxy by inferring the past nutritional
468 strategies of extinct taxa. Finally, the compositional differences between the two groups
469 suggest that preservation potential is influenced by dinosporin composition.

470

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479

480 REFERENCES

- 481 Allard, B., Templier, J. & Largeau, C. 1997. Artifactual origin of mycobacterial bacteran.
482 Formation of melanoidin-like artifactual macromolecular material during the usual isolation
483 process. *Org. Geochem.* 26: 691-703.
- 484
- 485 Allard, B., Templier, J. & Largeau, C. 1998. An improved method for the isolation of artifact-
486 free algaenans from microalgae. *Org. Geochem.* 28: 543-548.
- 487
- 488 Almendros, G., Dorado, J., González-Villa, F. J. & Martin, F. 1997. Pyrolysis of
489 carbohydrate-derived macromolecules: its potential in monitoring the carbohydrate signature
490 of geopolymers. *J. Anal. Appl. Pyrolysis* 40-41: 599-610.
- 491
- 492 Arnosti, C. 1995. Measurement of depth- and site-related differences in polysaccharide
493 hydrolysis rates in marine sediments. *Geochim. Cosmochim. Acta* 59: 4247-4257.
- 494
- 495 Aspinall, G. O. 1983. The polysaccharides. In Priess, J. [Ed.] *The Biochemistry of Plants*.
496 Academic Press, New York, NY, pp. 473-500.
- 497
- 498 Barker, S. A., Bourne, E. J., Stacey, M. & Whiffen, D. H. 1954. Infra-red spectra of
499 carbohydrates. Part I. Some derivatives of _D-glucopyranose. *J. Chem. Soc.* 171-176.
- 500
- 501 Bauersachs, T., Speelman, E. N., Hopmans, E. C., Reichart, G.-J., Schouten, S. & Sinninghe
502 Damsté, J. S. 2010. Fossilized glycolipids reveal past oceanic N₂ fixation by heterocystous
503 cyanobacteria. *Proc. Natl. Acad. Sci.* 107: 19190-19194.
- 504

505 Brenner, W. W. & Biebow, N. 2001. Missing autofluorescence of recent and fossil
506 dinoflagellate cysts – an indication of heterotrophy? *Neues Jahrb. Mineral. Geol. Palaeontol.*
507 *Abh. Abt. B.* 219: 229-240.

508

509 Bogus, K., Harding, I. C., King, A., Charles, A. K., Zonneveld, K. & Versteegh, G. J. M.
510 2012. The composition of species of the *Apectodinium* complex (Dinoflagellata). *Rev. of*
511 *Palaeobot. Palynol.* 183: 21-31.

512

513 Bouimetarhan, I., Marret, F., Dupont, L. & Zonneveld, K. 2009. Dinoflagellate cyst
514 distribution in marine surface sediments off West Africa (17-6°N) in relation to sea-surface
515 conditions, freshwater input and seasonal coastal upwelling. *Mar. Micropaleontol.* 71: 113-
516 130.

517

518 Bringué, M., Pospelova, V. & Pak, D. 2013. Seasonal production of organic-walled
519 dinoflagellate cysts in an upwelling system: A sediment trap study from the Santa Barbara
520 Basin, California. *Mar. Micropaleontol.* 100: 34-51.

521

522 Burton, R. A. & Fincher, G. B. 2009. (1,3; 1,4)- β -D-glucans in cell walls of the Poacea, lower
523 plants and fungi: a tale of two linkages. *Mol. Plant* 2: 873-882.

524

525 Cárdenas, G., Cabrera, G., Taboada, E. & Miranda, S. P. 2004. Chitin characterization by
526 SEM, FTIR, XRD, and ¹³C cross polarization/mass angle spinning NMR. *J. Appl. Polym.*
527 *Sci.* 93: 1876-1885.

528

529 Chapman, D. V., Dodge, J. D. & Heaney, S. I. 1982. Cyst formation in the freshwater
530 dinoflagellate *Ceratium hirundinella* (Dinophyceae). *J. Phycol.* 18: 121-129.

531

532 Coates, J. 2000. Interpretation of Infrared Spectra, A Practical Approach. In Meyers, R.A.
533 [Ed.] *Encyclopedia of Analytical Chemistry*. John Wiley and Sons Ltd., Chichester, pp.
534 10815-10837.

535

536 Colthup, N. B., Daly, L. H. & Wiberly, S. E. 1990. *Introduction to Infrared and Raman*
537 *Spectroscopy*. Academic Press Limited, London, 282 pp.

538

539 Combourieu-Nebout, N., Paterne, M., Turon, J. L. & Siani, G. 1998. A high-resolution record
540 of the last deglaciation in the central Mediterranean Sea: Palaeovegetation and
541 palaeohydrological evolution. *Quat. Sci. Rev.* 17: 303-317.

542

543 Dale, B. 1976. Cyst formation, sedimentation, and preservation: factors affecting
544 dinoflagellate assemblages in recent sediments from Trondheimsfjord, Norway. *Rev.*
545 *Palaeobot. Palynol.* 22: 39-60.

546

547 Dale, B., Dale, A. L. & Jansen, J. H. F. 2002. Dinoflagellate cysts as environmental indicators
548 in surface sediments from the Congo deep-sea fan and adjacent regions. *Palaeogeogr.*
549 *Palaeoclimatol. Palaeoecol.* 185: 309-338.

550

551 de Leeuw, J. W. 2007. On the origin of sedimentary aliphatic macromolecules: a comment on
552 recent publications by Gupta et al. *Org. Geochem.* 38: 1585-1587.

553

554 de Leeuw, J. W. & Largeau, C. 1993. A review of macromolecular organic compounds that
555 comprise living organisms and their role in kerogen, coal and petroleum formation. In Engel,

556 M. H. & Macko, S. A. [Eds.] *Organic Geochemistry: principles and applications*. Plenum
557 Publishing Corp., New York. pp. 23-72.
558

559 de Leeuw, J. W., Versteegh, G. J. M. & van Bergen, P. F. 2006. Biomacromolecules of algae
560 and plants and their fossil analogues. *Plant Ecol.* 182: 209-233.
561

562 Elbrächter, M. 1993. *Kolkwitziella* Lindemann 1919 and *Preperidinium* Mangin 1913: correct
563 genera names in the *Diplopsalis*-group (Dinophyceae). *Nova Hedwigia* 56: 173-178.
564

565 Ellegaard, M. 2000. Variations in dinoflagellate cyst morphology under conditions of
566 changing salinity during the last 2000 years in the Limfjord, Denmark. *Rev. Palaeobot.*
567 *Palynol.* 109: 65-81.
568

569 Ellegaard, M., Lewis, J. & Harding, I. 2002. Cyst-theca relationship, life cycle, and effects of
570 temperature and salinity on the cyst morphology of *Gonyaulax baltica* sp. nov. (Dinophyceae)
571 from the Baltic Sea area. *J. Phycol.* 38: 775-789.
572

573 Ellegaard, M., Figueroa, R. L. & Versteegh, G. J. M. 2013. Dinoflagellate life cycles, strategy
574 and diversity: key foci for future research. In Lewis, J. M., Marret, F. & Bradley, L. [Eds.]
575 *Biological and Geological Perspectives of Dinoflagellates*. The Micropalaeotological Society
576 Special Publications, Geological Society, London. pp. 249-261.
577

578 Fensome, R. A., Taylor, F. J. R., Norris, G., Sarjeant, W. A. S., Wharton, D. I. & Williams,
579 G. L. 1993. A classification of fossil and living dinoflagellates. *Micropaleontology Press*
580 *Special Paper*, 7, pp. 351.
581

582 Frei, E. & Preston, R. D. 1964. Non-cellulosic structural polysaccharides in algal cell walls I.
583 Xylan in siphonous green algae. *Proc. R. Soc. London, Ser. B.* 160: 293-313.
584

585 Fuentes-Grünewald, C., Garcés, E., Rossi, S. & Camp, J. 2009. Use of the dinoflagellate
586 *Karlodinium veneficum* as a sustainable source of biodiesel production. *J. Ind. Microbiol.*
587 *Biotechnol.* 36: 1215-1224.
588

589 Fuentes-Grünewald, C., Garcés, E., Alacid, E., Sampedro, N., Rossi, S. & Camp, J. 2012.
590 Improvement in lipid production in the marine strains *Alexandrium minutum* and
591 *Heterosigma akashiwo* by utilizing abiotic parameters. *J. Ind. Microbiol. Biotechnol.* 39: 207-
592 216.
593

594 Furuhashi, T., Beran, A., Blazso, M., Czegeny, Z., Schwarzinger, C. & Steiner, G. 2009.
595 Pyrolysis GC/MS and IR spectroscopy in chitin analysis of molluscan shells. *Biosci.*
596 *Biotechnol. Biochem.* 73: 93-103.
597

598 Gómez, F. 2012. A quantitative review of the lifestyle, habitat and trophic diversity of
599 dinoflagellates (Dinoflagellata, Alveolata). *Syst. Biodiversity* 10, 267-275.
600

601 González, C., Dupont, L. M., Mertens, K. & Wefer, G. 2008. Reconstructing marine
602 productivity of the Cariaco Basin during marine isotope stages 3 and 4 using organic-walled
603 dinoflagellate cysts. *Paleoceanography* 23, PA3215, doi:10.1029/2008PA001602.
604

605 Gupta, N. S., Collinson, M. E., Briggs, D. E. G., Evershed, R. P. & Pancost, R. 2006.
606 Reinvestigation of the occurrence of cutan in plants: implications for the leaf fossil record.
607 *Paleobiol.* 32: 432-449.

608

609 Hallet, R. I. 1999. Consequences of environmental change on the growth and morphology of
610 *Lingulodinium polyedrum* (Dinophyceae) in culture. Ph.D. dissertation, University of
611 Westminster, London, 109 pp.

612

613 Head, M. J. 1996. Modern dinoflagellate cysts and their biological affinities. *In* Jansonius, J.
614 & McGregor, D. C. [Eds.] *Palynology: Principles and Applications*. AASP Foundation, Salt
615 Lake City, UT, pp. 1197-1248.

616

617 Hemsley, A. R., Barrie, P. J., Scott, A. C. & Chaloner, W.G. 1994. Studies of fossil and
618 modern spore and pollen wall biomacromolecules using ^{13}C solid state NMR. *In* Eglinton, G.
619 & Kay, R. L. F. [Eds.] *Biomolecular Palaeontology*, NERC Special Publications, 94: 15-19.

620

621 Holzwarth, U., Esper, O. & Zonneveld, K. 2007. Distribution of organic-walled dinoflagellate
622 cysts in sediments of the Benguela upwelling system in relationship to environmental
623 conditions. *Mar. Micropaleontol.* 64: 91-119.

624

625 Holzwarth, U., Esper, O., Zonneveld, K. A. F. 2010. Organic-walled dinoflagellate cysts as
626 indicators of oceanographic conditions and terrigenous input in the NW African upwelling
627 region. *Rev. Palaeobot. Palynol.* 159: 35-55.

628

629 Hopkins, J. A. & McCarthy, F. M. G. 2002. Post-depositional palynomorph degradation in
630 Quaternary shelf sediments: a laboratory experiment studying the effects of progressive
631 oxidation. *Palynol.* 26: 167-184.

632

633 Hoppenrath, M. & Leander, B. S. 2010. Dinoflagellate Phylogeny as inferred from heat shock
634 protein 90 and ribosomal gene sequences. *PLoS ONE* 5: e13220.
635

636 Jacobson, D. M. & Anderson, D. M. 1986. Thecate heterotrophic dinoflagellates: feeding
637 behavior and mechanisms. *J. Phycol.* 22: 249-258.
638

639 Kačuráková, M. & Wilson, R.H. 2001. Developments in mid-infrared FT-IR spectroscopy of
640 selected carbohydrates. *Carbohydr. Polym.* 44: 291-303.
641

642 Kačuráková, M., Wellner, N., Ebringerová, A., Hromádková, Z., Wilson, R. H. & Belton, P.
643 S. 1999. Characterisation of xylan-type polysaccharides and associated cell wall components
644 by FT-IR and FT-Raman spectroscopies. *Food Hydrocolloids* 13: 35-41.
645

646 Kok, M. D., Schouten, S. & Sinninghe Damsté, J. S. 2000. Formation of insoluble,
647 nonhydrolyzable, sulfur-rich macromolecules via incorporation of inorganic sulfur species
648 into algal carbohydrates. *Geochim. Cosmochim. Acta* 64: 2689-2699.
649

650 Kokinos, J. P. & Anderson, D. M. 1995. Morphological development of resting cysts in
651 culture of the marine dinoflagellate *Lingulodinium polyedrum* (= *L. machaerophorum*).
652 *Palynol.* 19: 143-165.
653

654 Kokinos, J. P., Eglinton, T. I., Goñi, M. A., Boon, J. J., Martoglio P. A. & Anderson, D. M.
655 1998. Characterization of a highly resistant biomacromolecular material in the cell wall of a
656 marine dinoflagellate resting cyst. *Org. Geochem.* 28: 265-288.
657

658 Kouli, K., Brinkhuis, H. & Dale, B. 2001. *Spiniferites cruciformis*: a fresh water
659 dinoflagellate cyst? *Rev. Palaeobot. Palynol.* 113: 273-286.
660

661 Kurita, K. 2006. Chitin and chitosan: Functional biopolymers from marine crustaceans. *Mar.*
662 *Biotechnol.* 8: 203–226.
663

664 Lechien, V., Rodriguez, C., Ongena, M., Hilgsmann, S., Rulmont, A. & Thonart, P. 2006.
665 Physiochemical and biochemical characterization of non-biodegradable cellulose in Miocene
666 gymnosperm wood from the Entre-Sambre-et-Meuse, Southern Belgium. *Org. Geochem.* 37:
667 1465-1476.
668

669 Leroy, S. A. G., Boyraz, S. & Gürbüz, A. 2009. High-resolution palynological analysis in
670 Lake Sapanca as a tool to detect recent earthquakes on the North Anatolian Fault. *Quat. Sci.*
671 *Rev.* 28: 2616-2632.
672

673 Lessard, E. J. & Swift, E. 1986. Dinoflagellates from the North Atlantic classified as
674 phototrophic or heterotrophic by epifluorescence microscopy. *J. Plankton Res.* 8: 1209-1215.
675

676 Lewis, J., Rochon, A., Ellegaard, M., Mudie, P. J. & Harding, I. C. 2001. The cyst-theca
677 relationship of *Bitectatodinium tepikiense* (Dinophyceae). *Eur. J. Phycol.* 36: 137-146.
678

679 Marchand, C., Disnar, J. R., Lallier-Vergès, E. & Lottier, N. 2005. Early diagenesis of
680 carbohydrates and lignin in mangrove sediments subject to variable redox conditions (French
681 Guiana). *Geochim. Cosmochim. Acta* 69: 131-142.
682

683 Matsuoka, K. 1985. Cyst and thecate forms of *Pyrophacus steinii* (Schiller) Wall et Dale,
684 1971. *Transactions and proceedings of the Palaeontological Society of Japan, New series.*
685 140: 240-262.

686

687 Matsuoka, K. 1988. Cyst-theca relationships in the Diplopsalid group (Peridinales,
688 Dinophyceae). *Rev. Palaeobot. Palynol.* 56: 95-122.

689

690 Matsuoka, K., McMinn, A. & Wrenn, J. H. 1997. Restudy of the holotype of *Operculodinium*
691 *centrocarpum* (Deflandre & Cookson) Wall (Dinophyceae) from the Miocene of Australia,
692 and the taxonomy of related species. *Palynol.* 21: 19-33.

693

694 Matsuoka, K., Cho, H.-J. & Jacobson, D. M. 2000. Observations of the feeding behavior and
695 growth rates of the heterotrophic dinoflagellate *Polykrikos kofoidii* (Polykrikaceae,
696 Dinophyceae). *Phycologia* 39: 82-86.

697

698 Matsuoka, K., Kawami, H., Nagai, S., Iwataki, M. & Takayama, H. 2009. Re-examination of
699 cyst-motile relationships of *Polykrikos kofoidii* Chatton and *Polykrikos schwartzii* Bütschli
700 (Gymnodiniales, Dinophyceae). *Rev. Palaeobot. Palynol.* 154: 79-90.

701

702 Matthiessen, J., de Vernal, A., Head, M., Okolodkov, Y., Zonneveld, K. A. F. & Harland, R.
703 2005. Modern organic-walled dinoflagellate cysts in Arctic marine environments and their
704 (paleo-) environmental significance. *Palaeontolog. Z.* 79: 3-51.

705

706 McCarthy, F. M. G., Gostlin, K. E., Mudie, P. J. & Scott, D.B. 2000. Synchronous
707 palynological changes in early Pleistocene sediments off New Jersey and Iberia, and a
708 possible paleoceanographic explanation. *Palynol.* 24: 63-77.

709

710 McCarthy, F. M. G, Mertens, K. N., Ellegaard, M., Sherman, K., Pospelova, V., Ribeiro, S.,
711 Blasco, S. & Vercauteren, D. 2011. Resting cysts of freshwater dinoflagellates in southeastern
712 Georgian Bay (Lake Huron) as proxies of cultural eutrophication. *Rev. Palaeobot. Palynol.*
713 166: 46-62.

714

715 Menden-Deuer, S., Lessard, E. J., Satterberg, J. & Grünbaum, D. 2005. Growth and starvation
716 survival capacity of three species of the pallium feeding thecate dinoflagellate genus
717 *Protoperidinium* (Peridiniacea, Dinophyceae) distributions. *Aquat. Microbial Ecol.* 41: 145-
718 152.

719

720 Mertens, K. N., Verhoeven, K., Verleye, T., Louwye, S., Amorim, A., Ribeiro, S., Deaf, A.
721 S., Harding, I. C., De Schepper, S., González, C., Kodrans-Nsiah, M., de Vernal, A., Henry,
722 M., Radi, T., Dybkjaer, K., Poulsen, N. E., Feist-Burkhardt, S., Chitolie, J., Heilmann-
723 Clausen, C., Londeix, L., Turon, J.-L., Marret, F., Matthiessen, J., McCarthy, F. M. G.,
724 Prasad, V., Pospelova, V., Kyffin Hughes, J. E., Riding, J. B., Rochon, A., Sangiorgi, F.,
725 Welters, N., Sinclair, N., Thun, C., Soliman, A., van Nieuwenhove, N., Vink A. & Young, M.
726 2009a. Determining the absolute abundance of dinoflagellate cysts in recent marine
727 sediments: The *Lycopodium* marker-grain method put to the test. *Rev. Palaeobot. Palynol.*
728 157: 238-252.

729

730 Mertens, K. N., Ribeiro, S., Bouimetarhan, I., Caner, H., Combourieu-Nebout, N., Dale, B.,
731 de Vernal, A., Ellegaard, M., Filipova, M., Godhe, A., Goubert, E., Grøsfjeld, K., Holzwarth,
732 U., Kotthoff, U., Leroy, S. A. G., Londeix, L., Marret, F., Matsuoka, K., Mudie, P. J., Naudts,
733 L., Peña-Manjarrez, J. L., Persson, A., Popescu, S. M., Pospelova, V., Sangiorgi, F., van der
734 Meer, M. T. J., Vink A., Zonneveld, K. A. F., Vercauteren, D., Vlassenbroeck, J. & Louwye,

735 S., 2009b. Process length variation in cysts of a dinoflagellate, *Lingulodinium*
736 *machaerophorum*, in surface sediments: investigating its potential as salinity proxy. *Mar.*
737 *Micropal.* 70: 54-69.

738

739 Mertens, K. N., Rengefors, K., Moestrup, Ø. & Ellegaard, M. 2012. A review of recent
740 freshwater dinoflagellate cysts: taxonomy, phylogeny, ecology and palaeocology. *Phycologia*
741 51: 612-619.

742

743 Miller, A. A. L., Mudie, P. J. & Scott, D. B. 1982. Holocene history of Bedford Basin, Nova
744 Scotia: foraminifera, dinoflagellate and pollen records. *Can. J. Earth Sci.* 19: 2342-2367.

745

746 Moers, M. E. C., de Leeuw, J. W. & Baas, M. 1994. Origin and diagenesis of carbohydrates
747 in ancient sediments. *Org. Geochem.* 21: 1093-1106.

748

749 Naustvoll, L.-J. 2000. Prey size spectra and food preferences in thecate heterotrophic
750 dinoflagellates. *Phycologia* 39: 187-198.

751

752 Nevo, Z. & Sharon, N. 1969. The cell wall of *Peridinium westii*, a non cellulose glucan.
753 *Biochim. Biophys. Acta* 173: 161-175.

754

755 Orr, R. J. S., Murray, S. A., Stüken, A., Rhodes, L. & Jakobsen, K. S. 2012. When naked
756 became armored: An eight-gene phylogeny reveals monophyletic origin of theca in
757 dinoflagellates. *PLoS ONE* 7: e50004.

758

759 Pandey, K. K. 1999. A study of chemical structure of soft and hardwood and wood polymers
760 by FTIR spectroscopy. *J. Appl. Polym. Sci.* 71: 1969-1975.

761

762 Pfiester, L. A. & Anderson, D. M. 1987. Dinoflagellate reproduction. *In* Taylor, F. J. R. [Ed.]
763 *The Biology of Dinoflagellates*. Blackwell Scientific, Oxford, pp. 611-648.

764

765 Pospelova, V., Pedersen, T. F. & de Vernal, A. 2006. Dinoflagellate cysts as indicators of
766 climatic and oceanographic changes during the past 40 kyr in the Santa Barbara Basin,
767 southern California. *Paleoceanography* 21, PA2010, doi: 10.1029/2005PA001251.

768

769 Pospelova, V., de Vernal, A. & Pedersen, T. F. 2008. Distribution of dinoflagellate cysts in
770 surface sediments from the northeastern Pacific Ocean (43-25°N) in relation to sea-surface
771 temperature, salinity, productivity and coastal upwelling. *Mar. Micropal.* 68: 21-48.

772

773 Reichart, G. J. & Brinkhuis, H. 2003. Late Quaternary *Protoperidinium* cysts as indicators of
774 paleoproductivity in the northern Arabian Sea. *Mar. Micropal.* 49: 303-315.

775

776 Rochon, A., de Vernal, A., Turon, J. L., Matthiessen, J. & Head, M. J. 1999. Distribution of
777 recent dinoflagellate cysts in surface sediments from the North Atlantic Ocean and adjacent
778 seas in relation to sea-surface parameters. *American Association of Stratigraphic*
779 *Palynologists Foundation, Contribution series 35*, Dallas, TX, pp. 152.

780

781 Šandula, J., Kogan, G., Kačuráková, M. & Machová, E. 1999. Microbial (1-3)- β -D-glucans,
782 their preparation, physico-chemical characterization and immunomodulatory activity.
783 *Carbohydr. Polym.* 38: 247-253.

784

785 Schnepf, E. & Elbrächter, M. 1992. Nutritional strategies in dinoflagellates: A review with
786 emphasis on cell biological aspects. *Eur. J. Protistol.* 28: 3-24.

787

788 Sekida, S., Horiguchi, T. & Okuda, K. 2004. Development of thecal plates and pellicle in the
789 dinoflagellate *Scrippsiella hexapraecingula* (Peridinales, Dinophyceae) elucidated by
790 changes in stainability of the associated membranes. *Eur. J. Phycol.* 39: 105-114.

791

792 Stankiewicz, B. A., Mastalerz, M., Hof, C. H. J., Bierstedt, A., Flannery, M. B., Briggs, D. E.
793 G. & Evershed, R. P. 1998. Biodegradation of the chitin-protein complex in crustacean
794 cuticle. *Org. Geochem.* 28: 67-76.

795

796 Stone, B. A. 2009. Chemistry of β -glucans. In Bacic, A., Fincher, G. B. & Stone, B. A. [Eds.]
797 *Chemistry, biochemistry, and biology of (1-3)- β -glucans and related polysaccharides*.
798 Academic Press, Elsevier Inc., London, pp. 5-46.

799

800 Tardio, M., Sangiorgi, F., Ellegaard, M., Di Giuseppe, G., Filippi, M.L., Cantonati, M. &
801 Lotter, A. F. 2006. Peridinioid dinoflagellate cysts in a Holocene high-mountain lake deposits
802 in Italy. *J. Paleolimnol.* 36: 315-318.

803

804 Taylor, F. J. R., 1987. Ecology of dinoflagellates: general and marine ecosystems. In Taylor,
805 F. J. R. [Ed.] *The Biology of Dinoflagellates*. Botanical Monographs 21, Oxford, pp. 398-502.

806

807 Taylor, F.J.R. 2004. Illumination or confusion? Dinoflagellate molecular phylogenetic data
808 viewed from a primarily morphological standpoint. *Phycol. Res.* 52: 308-324.

809

810 Thornton, D. C. O., Santillo, D. & Thake, B. 1999. Prediction of sporadic mucilaginous algal
811 blooms in the northern Adriatic Sea. *Mar. Poll. Bull.* 38: 891-898.

812

813 van Dongen, B. E., Schouten, S., Baas, M., Geenevasen, J. A. J. & Sinninghe Damsté, J. S.
814 2003. An experimental study of the low-temperature sulfurization of carbohydrates. *Org.*
815 *Geochem.* 34: 1129-1144.

816

817 Venyaminov, S. Yu. & Kalnin, N.N. 1990. Quantitative IF spectrophotometry of peptide
818 compounds in water (H₂O) solutions, I. Spectral parameters of amino acid residue absorption
819 bands. *Biopolym.* 30: 1243-1257.

820

821 Verleye, T. J., Pospelova, V., Mertens, K. N. & Louwye, S. 2011. The geographical
822 distribution and (palaeo) ecology of *Selenopemphix undulata* sp. nov., a late Quaternary
823 dinoflagellate cyst from the Pacific Ocean. *Mar. Micropal.* 78: 65-83.

824

825 Versteegh, G. J. M., Blokker, P., Wood, G., Collinson, M. E., Sinninghe Damsté, J. S. & de
826 Leeuw, J. W. 2004. An example of oxidative polymerization of unsaturated fatty acids as a
827 preservation pathway for dinoflagellate organic matter. *Org. Geochem.* 35: 1129-1139.

828

829 Versteegh, G. J. M., Blokker, P., Marshall, C. P. & Pross, J. 2007. Macromolecular
830 composition of the dinoflagellate cyst *Thalassiphora pelagica* (Oligocene, SW Germany).
831 *Org. Geochem.* 38: 1643-1656.

832

833 Versteegh, G. J. M., Zonneveld, K. & de Lange, G. J. 2010. Selective aerobic and anaerobic
834 degradation of lipids and palynomorphs in the Eastern Mediterranean since the onset of
835 sapropel S1 deposition. *Mar. Geol.* 278: 177-192.

836

837 Versteegh, G. J. M., Blokker, P., Bogus, K., Harding, I. C., Lewis, J., Oltmanns, S., Rochon,
838 A. & Zonneveld, K. A. F. 2012. Flash pyrolysis and infrared spectroscopy of cultured and

839 sediment derived *Lingulodinium polyedrum* (Dinoflagellata) cyst walls. *Org. Geochem.* 43:
840 92-102.

841

842 Wall, D. & Dale, B. 1968. Modern dinoflagellate cysts and evolution of the Peridinales.
843 *Micropaleontol.* 14: 265–304.

844

845 Wotton, R. S. 2004. The ubiquity and many roles of exopolymers (EPS) in aquatic systems.
846 *Sci. Mar.* 68 (Suppl.): 13-21.

847

848 Yuen, S. N., Choi, S.-M., Phillips, D. E. & Ma, C.-Y. 2009. Raman and FTIR spectroscopic
849 study of carboxymethylated non-starch polysaccharides. *Food Chem.* 114: 1091-1098.

850

851 Zhang, H., Bhattacharya, D. & Lin, S. 2007. A three-gene dinoflagellate phylogeny suggests
852 monophyly of Prorocentrales and a basal position for Amphidinium and Heterocapsa. *J. Mol.*
853 *Evol.* 65: 463-474.

854

855 Zippi, P., Yung, Y.-K., McAndrews, J. A., Norris, G. & Welbourn, P. 1990. An investigation
856 of the potential of zygneatacean zygospores, Peridinium, and Pediastrum as paleo-
857 indicators of recent lake acidification. *Environmental Research, Technology Transfer*
858 *Conference (Toronto, Canada) Proceedings*, 1, pp. 393-396.

859

860 Zonneveld, K. A. F. & Susek, E. 2007. Effects of temperature, light and salinity on cyst
861 production and morphology of *Tuberculodinium vancampoae* (the resting cyst of *Pyrophacus*
862 *steinii*). *Rev. Palaeobot. Palynol.* 145: 77-88.

863

864 Zonneveld, K. A. F., Versteegh, G. J. M., & de Lange, G.J. 1997. Preservation of organic-
865 walled dinoflagellate cysts in different oxygen regimes: a 10,000 year natural experiment.
866 *Mar. Micropaleontol.* 29: 393-405.

867

868 Zonneveld, K. A. F., Versteegh, G. J. M. & de Lange, G. J. 2001. Palaeoproductivity and
869 post-depositional aerobic organic matter decay reflected by dinoflagellate cyst assemblages of
870 the Eastern Mediterranean S1 sapropel. *Mar. Geol.* 172: 181-195.

871

872 Zonneveld, K. A. F., Marret, F., Versteegh, G. J. M., Bogus, K., Bonnet, S., Bouimetarhan, I.,
873 Crouch, E., de Vernal, A., Elshanawany, R., Edwards, L., Esper, O., Forke, S., Grøsfjeld, K.,
874 Henry, M., Holzwarth, U., Kieft, J.-F., Kim, S.-Y., Ladouceur, S., Ledu, D., Chen, L.,
875 Limoges, A., Londeix, L., Lu, S.-H., Mahmoud, M., Marino, G., Matsuoka, K., Matthiessen,
876 J., Mildenhall, D. C., Mudie, P., Neil, H. L., Pospelova, V., Qi, Y., Radi, T., Richerol, T.,
877 Rochon, A., Sangiorgi, F., Solignac, S., Turon, J.-L., Verleye, T., Wang, Y. & Young, M.
878 2013. Atlas of modern dinoflagellate cyst distribution based on 2405 data points. *Rev.*
879 *Palaeobot. Palynol.* 191: 1-197.

880

881 *Tables*

882 Table 1: Surface sediment and culture sample information from which the dinocyst specimens

883 were isolated.

#	Name	Location	Latitude	Longitude	Water depth (m)	Setting	Reference
1	GeoB 2341	Benguela upwelling	31°55'48"S	18°12'36"E	84	Marine	Holzwarth et al. (2007)
2	GeoB 4804	Benguela upwelling	24°8'60"S	12°40'12"E	2090	Marine	Holzwarth et al. (2007)
3	GeoB 6010	NE Atlantic (off NW Africa)	30°15'N	10°2'16.8"W	406	Marine	Holzwarth et al. (2010)
4	SV5-C	Honey Harbour (Lake Huron, Canada)	44°52'26"N	79°48'55"W	19.4	Lacustrine	McCarthy et al. (2011)
5	Schillig	Wadden Sea (NW Germany)	53°42'56"N	7°58'11.93"E	0.5	Culture	-
6	GeoB 1010	Omura Bay (Kyushu, Japan)	33°N	128°49'48"E	Not stated	Culture	Zonneveld and Susek (2007)

884

885

886 Table 2: Dinocyst species analyzed in this study, their motile affinities, and nutritional
 887 strategy with reference to the respective studies.

Sample*	Specimens measured (#)	Dinocyst species	Motile affinity	Order
2	3	<i>Impagidinium patulum</i> (Wall) Stover and Evitt	<i>Gonyaulax</i> sp. Diesing ^a	Gonyaulacales
2,3	6 (3 from each location)	<i>Operculodinium centrocarpum</i> sensu Wall and Dale	<i>Protoceratium reticulatum</i> (Claparède et Lachmann) Bütschli ^b	Gonyaulacales
2	4	<i>Spiniferites pachydermus</i> (Rossignol) Reid	<i>Gonyaulax</i> sp. Diesing ^c	Gonyaulacales
6	3	<i>Tuberculodinium vancampoae</i> (Rossignol) Wall	<i>Pyrophacus steinii</i> (Schiller) Wall and Dale ^d	Gonyaulacales
4	3	Cyst of <i>Peridinium wisconsinense</i> Eddy	<i>Peridinium wisconsinense</i> Eddy ^e	Peridinales
1	4	<i>Brigantedinium</i> spp. Reid	<i>Protoperidinium</i> sp. Bergh ^c	Peridinales
5	3	<i>Dubridinium caperatum</i> Reid	<i>Preperidinium meunieri</i> (Pavillard) Elbrächter ^f	Peridinales
1	4	Cyst of <i>Polykrikos kofoidii</i> Chatton	<i>Polykrikos kofoidii</i> Chatton ^g	Gymnodinales
1,2	6 (3 from each location)	Cyst of <i>Polykrikos schwartzii</i> Bütschli	<i>Polykrikos schwartzii</i> Bütschli ^g	Gymnodinales

888 *Sample numbers correspond to Table 1. ^aRochon et al. (1999). ^bMatsuoka et al. (1997).

889 ^cThis study. ^dMatsuoka (1985). ^eMcCarthy et al. (2011). ^fMatsuoka (1988); see Elbrächter
 890 (1993) for synonymy. ^gMatsuoka et al. (2009). ^hGómez (2012).

891

892

893 Table 3: Assignments of major FTIR absorptions present in Group I (phototrophic) and
 894 Group II (heterotrophic) dinosporins. See Table 2 for each species' nutritional strategy.

Dinocysts	Wavenumber (cm ⁻¹)	Assignment	Comments
<u>Group I</u>			
	3348	vOH	
	2925	vCH	
	2860	vCH	
	1640	vC=O	
	1600	vC=C + vC=O	
	1430	δCH ₂	
	1370	δCH + δC-CH ₃	
	1318	δOH	
	1163	vC-O-C	
	1112	vC-O	Glucose ring
	1059	vC-O	
	1033	vC-O	
	897	γCH	β-glycosidic bond
<u>Group II</u>			
	3479, 3448, 3426	vOH	
	3268	vNH	
	3106	vNH	
	2965-2860	vCH	
	1660, 1627	vC=O	Amide I
	1585-1550	vCN+ δNH	Amide II
	1420	δCH ₂	

	1370	$\delta\text{CH} + \delta\text{C-CH}_3$	
	1312	$\nu\text{CN} + \delta\text{NH}$	Amide III
	1255	δNH	
	1157	$\nu\text{C-O-C}$	Ring
	1113	$\nu\text{C-O}$	Glucose ring
	1060	$\nu\text{C-O}$	
	1027	$\nu\text{C-O}$	
	896, 902	γCH	β -glycosidic bond
	875-746	ρCH_2	
	698	γNH	Amide V

895

896

897 *Figure captions*

898 Figure 1: Locations of the marine and lacustrine surface samples analyzed in this study.

899

900 Figure 2: FTIR spectra of the Group I (phototrophic) dinocysts compared to *Lingulodinium*
901 *machaerophorum* from sediment (black line) and culture (gray line; Versteegh et al. 2012)
902 and cellulose (Pandey 1999). The gray *O. centrocarpum* spectrum denotes the analysis from
903 the second measured location (see Tables 1, 2). Absorption assignments are given in Table 3.

904

905 Figure 3: FTIR spectra of the Group II (heterotrophic) dinocysts and the chitin standard. The
906 gray line for cysts of *P. schwartzii* reflects spectral data from a second location (see Tables 1,
907 2). Absorption assignments are given in Table 3.

908

909 Figure 4: Phylogenetic subdivision of the major orders according to phylogenies proposed by
910 Fensome et al. (1993) and Taylor (2004) and comparison with the two groups derived from
911 micro-FTIR measurements.

912