

1 **Soil inorganic N leaching in edges of different forest types**
2 **subject to high N deposition loads**

3 Short title: N leaching in edges of different forest types

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14 **Abstract**

15 We report on soil leaching of dissolved inorganic nitrogen (DIN) along transects across exposed
16 edges of four coniferous and four deciduous forest stands. In a 64-m edge zone, DIN leaching
17 below the main rooting zone was enhanced relative to the interior (at 128 m from the edge) by 21
18 and 14 kg N ha⁻¹ y⁻¹ in the coniferous and deciduous forest stands, respectively. However, the
19 patterns of DIN leaching did not univocally reflect those of DIN throughfall deposition. DIN
20 leaching in the first 20 m of the edges was lower than at 32-64 m from the edge (17 vs. 36 and 15
21 vs. 24 kg N ha⁻¹ y⁻¹ in the coniferous and deciduous forests, respectively). Nitrogen stocks in the
22 mineral topsoil (0-30 cm) were, on average, 943 kg N ha⁻¹ higher at the outer edges than in the
23 interior, indicating that N retention in the soil is probably one of the processes involved in the
24 relatively low DIN leaching in the outer edges. We suggest that a complex of edge effects on
25 biogeochemical processes occurs at the forest edges as a result of the interaction between
26 microclimate, tree dynamics (growth and litterfall), and atmospheric deposition of N and base
27 cations.

28

29 **Keywords**

30 atmospheric deposition / nitrate leaching / forest edge / edge effects / forest type / C:N ratio /
31 inorganic nitrogen

32 1. Introduction

33 In many landscapes around the world, human activities such as timber harvesting, agricultural
34 expansion, and urbanization increase the extent of forest fragmentation, making the forest edge the
35 prevailing feature of the landscape matrix (Riitters and others 2002; Harper and others 2005;
36 Broadbent and others 2008). In forest edges, nitrogen (N) throughfall deposition is enhanced
37 relative to the forest interior (see De Schrijver and others (2007a) for a review). In the interior of N
38 saturated forests, high levels of N deposition are associated with increased levels of dissolved
39 inorganic N (DIN; ammonium (NH_4^+) + nitrate (NO_3^-)) leaching to the groundwater, predominantly
40 as NO_3^- (MacDonald and others 2002; Kristensen and others 2004; Vestgarden and others 2004; De
41 Schrijver and others 2008). Hence, higher levels of DIN leaching in forest edges in comparison with
42 the forest interior can be expected. Yet, both lower (Balsberg-Påhlsson and Bergkvist 1995;
43 Kinniburgh and Trafford 1996; Spangenberg and Kölling 2004) and higher (Balsberg-Påhlsson and
44 Bergkvist 1995; Spangenberg and Kölling 2004) NO_3^- concentrations in soil water and DIN
45 leaching fluxes have been detected at edges than further down the edge.

46 The proximity of the edge affects a variety of abiotic and biotic variables. Microclimatic variables
47 such as air and soil temperature, light availability, and soil moisture are altered (Matlack 1993;
48 Chen and others 1995; Gehlhausen and others 2000; Marchand and Houle 2006; Heithecker and
49 Halpern 2007) and throughfall deposition fluxes of sulphur and base cations such as K^+ , Ca^{2+} , and
50 Mg^{2+} are enhanced (Draaijers and others 1994; Spangenberg and Kölling 2004; Wuyts and others
51 2008a). Moreover, higher tree growth rates and leaf area indices are reported at edges in
52 comparison with forest interiors (McDonald and Urban 2004; Bowering and others 2006; Sherich
53 and others 2007), potentially resulting in higher levels of N sequestration in biomass at the edge due
54 to higher tree biomass (Spangenberg and Kölling 2004).

55 Our research objective was to investigate (i) whether DIN leaching in the soil is influenced by edge
56 effects in both deciduous and coniferous forest types, and (ii) how this pattern of DIN leaching in

57 edges relates to patterns of DIN throughfall deposition, of forest floor and mineral soil
58 characteristics, and of N stored in stem biomass. In addition, we examined the effect of forest type
59 on DIN leaching. We presume that the higher input of N deposition is expressed in enhanced DIN
60 leaching near edges, and that this pattern of DIN leaching differs between forest types, since
61 deciduous forest types are subject to smaller edge effects on DIN deposition than coniferous ones
62 are (Spangenberg and Kölling 2004; Wuyts and others 2008a, 2008b).

63

64 **2. Materials and methods**

65 2.1. Site description

66 In two regions, located in Flanders (the northern part of Belgium) and approximately 135 km apart,
67 eight homogeneous forest stands were considered (Fig. 1): two naturally regenerated stands
68 dominated by silver birch (*Betula pendula* Roth; stands Bp1 and Bp2) and six planted monocultures
69 of (i) pedunculate oak (*Quercus robur* L.; stands Qr1 and Qr2), (ii) Austrian pine (*P. nigra* ssp.
70 *nigra* Arnold; stand Pn1) or Corsican pine (*Pinus nigra* ssp. *laricio* Maire; stand Pn2), and (iii)
71 Scots pine (*Pinus sylvestris* L.; stands Ps1 and Ps2). The stands Bp1, Qr1, Pn1, and Ps1 were
72 situated in the western part of Flanders and Bp2, Qr2, Pn2, and Ps2 in the northern part (Fig. 1).
73 Site code, location, stand characteristics, and soil pH-KCl of the studied forest stands are presented
74 in Table 1. Thinning in stand Ps2 about three years prior to the start of our measuring campaign
75 caused DIN leaching to be locally and temporarily enhanced, higher than the input of N via
76 throughfall deposition (von Wilpert and others 2000; Weis and others 2006). Hence, DIN leaching
77 data from the Ps2 site was omitted from statistical analyses and figures.

78 All stands had an abrupt forest edge oriented towards the prevailing southwesterly winds. All
79 forests were located on poor, well-drained soils of sand or loamy sand (Haplic podzols; FAO-
80 ISRIC-ISSS 1998) with a low buffering capacity for acidifying deposition (Van Ranst and others
81 2002; De Schrijver and others 2006). During the entire sampling period, the groundwater table was

82 deeper than 150 cm in all stands except Pn1 (> 120 cm) and Bp1 and Qr1 (> 80 cm). All forests
83 were situated in areas with high density of intensive livestock farms. Analysis of the throughfall
84 deposition data (Wuyts and others 2008b) showed that differences in edge effects between the two
85 regions were less important than the differences between coniferous and deciduous forest stands.
86 For the physical and chemical characterisation of the climate in the two regions, we refer to Wuyts
87 and others (2008b). Historic maps from around 1775 and later showed that the sites were managed
88 as heath, characterised by sheep grazing and turf cutting, until up to 80 years ago, and, from then
89 onwards, were continuously covered by forest. The Qr1 stand was briefly managed as grassland
90 around 1900. The edge of the Pn1 stand borders a busy highway at less than 5 m.

91

92 2.2. Experimental setup and sampling

93 In each stand, throughfall and soil solution were monitored year-round, soil samples were taken,
94 and leaf area index (LAI) was determined at eight distance plots at 0, 2, 4, 8, 16, 32, 64, and 128 m
95 from the edge, along a transect perpendicular to the forest edge. Throughfall water was collected
96 using three collectors (funnels mounted on two-liter bottles) per distance plot (Wuyts and others
97 2008b). Soil solution was sampled at two depths, i.e., at 30 and 90 cm, by means of one suction cup
98 lysimeter per depth and distance plot. Within one plot, the distance between the two lysimeters was
99 at least 1 m. The lysimeters consisted of (i) a PVC pipe fitted with a porous ceramic cup (5 cm
100 diameter, Eijkelkamp, The Netherlands) and inserted into the soil at an angle of 60° with the surface
101 and (ii) an opaque, glass, one-liter bottle connected to the pipe via a polyethylene tube and stored
102 belowground to keep samples cool. After each sampling, a pressure of -500 hPa was applied.
103 Preferably, the lysimeters were installed where no understory vegetation occurred to avoid
104 confounding influence of its N uptake. Brambles (*Rubus fruticosus* agg.) were carefully removed
105 within a 3-m radius around the lysimeters in the first 20 m of Qr1, at the 64 m and 128 m distance
106 plots of Pn1, and at the 16 m and 32 m distance plots of Ps1. The perennial herb *Holcus mollis* L. in

107 the first 10 m from the edge of the Pn2 stand was not removed to avoid substantial disruption of the
108 organic layer. The lysimeters were installed four months prior to the actual monitoring to check for
109 proper performance, and were drained every two weeks to ensure close contact with the
110 surrounding soil solution. The level of groundwater was monitored with piezometers; in the stands
111 Qr1 and Bp1, soil solution was not sampled at 90 cm due to high ground water levels (> 80 cm
112 deep). Soil solution and throughfall sampling took place fortnightly from 1 September 2005 to 30
113 August 2006. On each sampling occasion, sample volume in the throughfall collectors and the
114 lysimeters was measured, throughfall samples were pooled volume-weighted per distance plot, and
115 subsamples were taken to the lab. Aliquots of two consecutive sample collections were pooled
116 volume-weighted into monthly samples for chemical analysis.

117 In February 2006, soil samples were taken with a 3.2 cm diameter soil auger from the fermentation
118 and humus layer (together, FH) of the forest floor (i.e., the ectorganic layer) and from 0-5 cm, 5-10
119 cm, and 10-30 cm of the mineral soil. We took samples at three locations per distance plot and all
120 samples were analysed separately. Next to this, the FH layer was collected in a 39 cm x 39 cm
121 square at the outer edge (0-20 m from the edge front) and the forest interior (at 128 m from the
122 edge). Density of the mineral soil was determined from the dry weight of undisturbed samples in
123 cylinders of approximately 10^{-4} m^3 .

124 In winter 2005-2006 and summer 2006, LAI was measured from digital hemispherical photographs
125 (Nikon D70s with fish-eye lens Sigma EXDG Fisheye 8 mm 1:4 D) taken above each throughfall
126 collector, using Gap Light Analyzer GLA 2.0 (<http://www.ecostudies.org/gla>). Stem volume was
127 calculated from diameter at breast height and tree height measured for all trees within 10 m from
128 the central transect axis in spring 2006, based on the equations by Jansen and others (1996).

129 In June 2008, leaves and current-year and older (> 1 year) needles were collected using telescopic
130 secateurs at a fixed height of 12 m and from the shaded part of the crown. We sampled five trees
131 located in the first 10 m of the forest edge and five trees in the forest interior. In the Qr1 stand,

132 branches with leaves were too high to reach, while in the Ps1 stand, needles could not be collected
133 due to safety reasons (next to a busy highway). Trees at the edges were sampled at the sheltered
134 side opposite from the edge to avoid overestimation of the N concentrations by sampling sunlit
135 leaves or needles. Dry weight and surface area of a leaf or needle subsample were measured to
136 assess the specific leaf area (SLA, ratio of leaf area to leaf dry weight). Leaf surface areas were
137 determined with the LI-3000 Portable Area Meter (LICOR); needle surface areas were estimated
138 from needle length and diameter. Stem cores were taken from the same trees at 30 cm height above
139 the root collar. Bark was excluded from chemical analysis and only fully intact stem cores, bored
140 through the entire stem, were selected. Additionally, the length, diameter, and weight of the dry
141 stem cores were determined to estimate wood density.

142

143 2.3. Chemical analyses

144 Subsequent to filtration (0.45 μm , Rothe), the monthly water samples were analyzed for NH_4^+
145 photometrically (determination of a reaction product of NH_4^+ at $\lambda = 660 \text{ nm}$ according to the Dutch
146 standard method NEN 6576; Cary 50 spectrophotometer, Varian) and for Cl^- and NO_3^- by ion
147 chromatography (ICS-90 Dionex). The forest floor samples were dried at 70°C and grinded;
148 samples of the mineral soil were dried at 40°C and sieved (2 mm mesh width). Samples of the FH
149 layer of the forest floor and the mineral soil were analysed for total N (salicylic acid-thiosulphate
150 modification of Kjeldahl method: 1 g soil or 0.2 g ectorganic material in 7 ml $\text{C}_7\text{H}_6\text{O}_3$ solution (50 g
151 dissolved in 1 l H_2SO_4 96%), 5 ml concentrated H_2SO_4 , and 0.5 g $\text{Na}_2\text{S}_2\text{O}_3 \cdot 5\text{H}_2\text{O}$; Bremner 1996)
152 and total C (estimated by loss on ignition: 4.5 h at increasing temperature until 450°C , then 4 h at
153 450°C ; estimated % C = $(100 - \% \text{ of ashes residue}) / 2$). The samples of the FH layer were digested
154 (0.2 g organic material in 10 ml HNO_3 and 2 ml HClO_4 ; McKenzie 2010) and the 0-5 cm samples
155 of the mineral soil were extracted with BaCl_2 (5.00 g soil in 100 ml 0.1 M BaCl_2 solution; Dutch
156 standard method NEN 5738) and were subsequently analysed for Ca^{2+} and Mg^{2+} (flame atomic

157 absorption spectrophotometry, SpectrAA-220, Varian). Wood and foliar material were dried at
158 70°C, grinded, and following total destruction analysed for total N. The quality of the chemical
159 analyses was checked by including method blanks and repeated measurements of internal standards
160 and certified reference samples.

161

162 2.4. Element fluxes, N stocks, and soil properties

163 Annual throughfall fluxes ($\text{kg ha}^{-1} \text{y}^{-1}$) were the sum of monthly fluxes calculated from the
164 concentration in the subsamples multiplied by the monthly throughfall volume in the three
165 collectors and divided by the surface area of the three funnels (Wuyts and others 2008b). Year-
166 round mean NO_3^- -N concentrations (mg l^{-1}) were calculated as the mean of NO_3^- -N concentrations
167 weighted by the sample volume in the lysimeters on each sampling event. Soil NH_4^+ -N and NO_3^- -N
168 leaching fluxes ($\text{kg ha}^{-1} \text{y}^{-1}$) were determined by multiplying the water seepage flux ($\text{l m}^{-2} \text{y}^{-1}$) with
169 the year-round, volume-weighted mean ion concentration in the soil solution (mg l^{-1}). The water
170 seepage fluxes were estimated with the chloride mass balance method (Eriksson and Khunakasem
171 1969; Boxman and others 2008; De Schrijver and others 2008). Recent studies indicate that Cl^-
172 release or retention may occur in the soil (Öberg and Sandén 2005; Bastviken and others 2007), but
173 we found the mass balance method with Cl^- as a tracer sufficient to study edge effects. The water
174 seepage fluxes in stands Qr2 and Pn2 calculated with Na^+ (flame atomic absorption
175 spectrophotometry, SpectrAA-220, Varian) as a tracer were lower than with Cl^- , but the difference
176 was marginally insignificant (paired samples t test, $P = 0.077$) and the Na^+ and Cl^- concentrations in
177 the soil solution were highly correlated irrespective of time and distance from the edge (Pearson R
178 $= 0.94$; $P < 0.001$; $n = 35$). The calculated water seepage fluxes were independent of edge distance.
179 However, at the first two distance plots at the Qr2 site, the calculated water seepage fluxes at 90 cm
180 soil depth were high and equalled the throughfall volumes ($\pm 600 \text{ l m}^{-2}$), while further down the
181 edge, they were half the volume of throughfall. A small but profound ($\pm 1 \text{ m}$) creek located 2 m

182 before the edge front of stand Qr2 may have influenced subsurface water flows. As a result, DIN
183 throughfall fluxes at the first two distance plots of Qr2 were probably overestimated.
184 The N stock in leaves or needles was approximated by multiplying the mean N concentration (kg N
185 kg^{-1}) by the LAI ($\text{m}^2 \text{m}^{-2}$, converted to $\text{m}^2 \text{ha}^{-1}$) and the inverse of the mean SLA. The N stock in
186 tree stem biomass was estimated by multiplying the mean N concentration (kg N kg^{-1}) by the stem
187 volume ($\text{m}^3 \text{ha}^{-1}$) and the wood density, which was in general lower at the edge than in the interior.
188 The C:N ratios were determined for the FH layer of the forest floor (C:N_{FH}) and the mineral soil
189 (C:N₀₋₅, C:N₅₋₁₀, C:N₁₀₋₃₀). The N stocks (kg N ha^{-1}) in the mineral soil were calculated based on the
190 measured N concentration and soil density (the latter was significantly lower at the edge than in the
191 interior).

192

193 2.5. Statistics

194 By means of a repeated-measures ANOVA (RMA), the effects of within-subject factor ‘distance to
195 the forest edge’ (8 levels) and between-subjects factor ‘forest type’ (2 levels: coniferous and
196 deciduous) on NO_3^- -N concentrations in soil water, dissolved inorganic nitrogen (DIN, NH_4^+ -N +
197 NO_3^- -N) leaching, and soil characteristics (C, N, C:N) were tested. Furthermore, we estimated the
198 mean NO_3^- -N concentrations and DIN soil leaching fluxes in the first 64 m of the edge, weighted by
199 the share of each distance plot in the considered zone (i.e., 1, 2, 3, 6, 12, 24, and 16 m for the first to
200 seventh distance plots, respectively). Repeated-measures ANOVA was applied to identify and
201 separate the influence of ‘edge proximity’ (2 levels: forest edge 0-64 m and forest interior at 128 m)
202 and forest type on NO_3^- -N concentrations and DIN leaching. This data pooling enabled to make
203 more general conclusions. Although the NH_4^+ throughfall deposition, representing on average 73%
204 of the DIN throughfall deposition, was significantly lower in region 1 than in region 2 by 7 kg N ha^{-1}
205 y^{-1} on average (Wuyts and others 2008b), we observed no significant influence of region and its

206 interaction effects on DIN leaching and N content. The factor region and its interaction effects were
207 therefore omitted from the analyses to decrease the number of parameter estimations.
208 Furthermore, differences between edge (0-64 m) and interior (128 m) were tested for the deciduous
209 and coniferous stands separately with non-parametric related samples tests (Wilcoxon signed ranks
210 test). A repeated-measures analysis (RMA) was applied on the dry weight of the FH layer, the N
211 stock in the mineral soil and the biomass (or LAI), the N concentration, and the N stock in stem
212 wood and leaf/needle to separate the influence of ‘edge proximity’ (2 levels: edge and interior) and
213 of forest type. Finally, single linear regression models (including constants) were tested between
214 DIN in throughfall deposition, $C:N_{FH}$, and the C:N ratios of the mineral soil on the one hand and the
215 (log-transformed) DIN leaching fluxes at 30 and 90 cm soil depth on the other hand. All statistical
216 analyses were performed using SPSS 15.0.

217

218 **3. Results**

219 3.1. DIN leaching

220 Higher NO_3^- -N concentrations in soil water occurred in the coniferous stands in comparison with
221 the deciduous ones. For example, in the interior part (128 m) of the studied forests, mean
222 concentrations \pm standard error were $14 \pm 3 \text{ mg l}^{-1}$ at 30 cm and $17 \pm 9 \text{ mg l}^{-1}$ at 90 cm soil depth in
223 the coniferous stands and $4.2 \pm 0.9 \text{ mg l}^{-1}$ at 30 cm and $3.3 \pm 0.3 \text{ mg l}^{-1}$ at 90 cm soil depth in the
224 deciduous stands. Similarly, estimated DIN leaching fluxes were higher in the coniferous than in the
225 deciduous stands, with mean interior fluxes of $35 \pm 7 \text{ kg N ha}^{-1} \text{ y}^{-1}$ at 30 cm and $10 \pm 5 \text{ kg N ha}^{-1} \text{ y}^{-1}$
226 at 90 cm soil depth in the coniferous stands and $8 \pm 2 \text{ kg N ha}^{-1} \text{ y}^{-1}$ at 30 cm and $7 \pm 2 \text{ kg N ha}^{-1} \text{ y}^{-1}$
227 at 90 cm soil depth in the deciduous stands. However, forest type affected DIN leaching only
228 significantly ($P < 0.05$) at 30 cm soil depth, and yet marginally insignificantly ($0.05 < P < 0.1$)
229 affected the NO_3^- -N concentrations at this depth (Table 2).

230 According to the RMA analysis (Table 2), the distance from the forest edge did not significantly (P
231 > 0.1) influence the NO₃⁻-N concentrations and the DIN leaching fluxes at 30 and 90 cm depth,
232 indicating that the NO₃⁻-N concentrations and the DIN leaching did not display a monotonic
233 (increasing or decreasing) relationship with distance from the edge. The NO₃⁻-N concentrations in
234 the soil water at 30 and 90 cm (Fig. 2, Appendix) were not enhanced at the edges, except for Pn1 at
235 30 cm depth, where we observed a slight decrease in concentration with increasing edge distance.
236 Instead, the NO₃⁻-N concentrations were highest between 8 and 32 m or increased towards the forest
237 interior (at both sampling depths in Pn2 and in Qr1 and Ps1 at 30 cm). In all stands, except for Pn1
238 at 30 cm and Qr2, the lowest DIN leaching fluxes occurred at the edge front (0 to 2 m) with values
239 equal or even lower than at 128 m, while maximum values were obtained at 8, 16, 32, or 64 m from
240 the edge (see Appendix). Fig. 2 shows the dissimilarity between the mean edge patterns of DIN
241 throughfall deposition and of DIN leaching fluxes in the deciduous and coniferous stands. While the
242 DIN throughfall deposition decreased with increasing edge distance, the DIN leaching fluxes at 90
243 cm soil depth increased with distance in the first 20 m of the edge ('outer edge'), peaked at 32 and
244 64 m, and decreased in the forest interior at 128 m. The DIN leaching at 90 cm soil depth in the first
245 20 m from the edge was 15 ± 12 and 17 ± 6 kg N ha⁻¹ y⁻¹ (mean \pm SE) in the deciduous and
246 coniferous forest stands, resp., while from 20 to 64 m the leaching amounted to 24 ± 9 and 36 ± 11
247 kg N ha⁻¹ y⁻¹, resp. No significant interaction between forest type and distance from the forest edge
248 occurred for NO₃⁻-N concentrations and DIN leaching at 30 and 90 cm (Table 2).
249 The NO₃⁻-N concentrations and leaching of DIN averaged over the first 64 m of the edge were
250 consistently higher than the concentrations and fluxes in the interior, and differences were
251 marginally insignificant for the leaching fluxes at 30 cm and significant for those at 90 cm (Table
252 2). The differences in NO₃⁻-N concentrations were, however, not significant. When considering
253 coniferous and deciduous stands separately, NO₃⁻-N concentrations and DIN leaching did not differ
254 significantly between the first 64 m of the edge and the interior, although the difference in DIN

255 leaching flux at 90 cm in the coniferous stands was only marginally insignificant. In the first 64 m
256 of the edges, the DIN leaching fluxes at 30 cm soil depth were in the deciduous stands, on average,
257 10 kg N ha⁻¹ y⁻¹ (125%) higher than in the forest interior and in the coniferous stands 24 kg N ha⁻¹ y⁻¹
258 (69%) higher than in the interior (Table 2). At 90 cm soil depth, the increase in DIN leaching in
259 the 64 m edge zone relative to the interior was 14 kg N ha⁻¹ y⁻¹ (or 181%) and 21 kg N ha⁻¹ y⁻¹ (or
260 105%) in the deciduous and coniferous stands, respectively (Table 2). For comparison, we provide
261 the results of the repeated measures analysis on DIN throughfall deposition and its weighted-means
262 in the 64-m edge zone and the interior of the deciduous and coniferous stands in Table 2.
263 Significant linear relations were found between NO₃⁻-N concentrations at 30 and 90 cm ($R^2 = 0.57$;
264 $P < 0.001$; slope = 1.982) and between the log-transformed DIN leaching fluxes at 30 and 90 cm
265 ($R^2 = 0.51$; $P < 0.001$; slope = 0.952). The DIN leaching fluxes at 30 cm and 90 cm were plotted
266 against the DIN in throughfall deposition in Fig.3. The DIN leaching fluxes at 30 cm were
267 significantly linearly related with the DIN input via throughfall deposition ($R^2 = 0.43$; $P < 0.001$;
268 slope = 1.51; Fig. 3). This relationship held when the deciduous stands were considered separately
269 ($R^2 = 0.33$; $P = 0.001$; slope = 1.79) but was marginally insignificant for the coniferous stands ($R^2 =$
270 0.14; $P = 0.073$; slope = 1.05). The linear relationship between DIN throughfall deposition and DIN
271 leaching was improved in the coniferous stands when only plots with $C:N_{FH} \leq 25$ (n=18) were
272 considered ($R^2 = 0.32$; $P = 0.014$; slope = 1.70). No significant linear relationships were observed
273 between the DIN throughfall deposition and the DIN leaching at 90 cm ($P = 0.138, 0.316, \text{ and } 0.377$
274 for all (n = 40), the deciduous (n = 16), and the coniferous (n = 24) stands, respectively).

275

276 3.2. Forest floor and mineral soil

277 At the outer edges (0-20 m), on average 7.5 ± 0.7 kg m⁻² of dry FH layer was collected, which was
278 significantly lower than in the forest interiors (10.9 ± 1.4 kg m⁻²; Table 3). A significant effect of
279 forest type was observed on the C concentrations and the C:N ratio of the FH layer (Table 4), which

280 were both higher in the deciduous than in the coniferous stands. The distance from the edge
281 significantly affected the C and N concentrations in the FH layer of the forest floor, as well as its
282 C:N ratio (Table 4). The overall trend was that the N and C concentrations and the C:N ratios were
283 lower at the edges (Fig. 4). No significant interactions between edge effects and forest type were
284 detected in the FH layer.

285 The N concentrations in the mineral soil were in general higher at the edges in all soil depths but
286 were significantly influenced by edge effects only in the 0-5 cm mineral topsoil and marginally
287 insignificantly at 10-30 cm (Table 4; Fig. 4). The C concentrations in the soil were significantly
288 affected by distance from the edge in all sampled soil depths (Table 4; Fig. 4), with higher
289 concentrations near the edge. The C:N ratios of the mineral topsoil (0-5 cm) were significantly
290 lower at the edge. The forest type had a significant effect on the N concentrations of the 0-5 cm and
291 5-10 cm mineral soil layers and the C:N ratios of the 0-5 cm mineral topsoil and had a marginally
292 insignificant effect on the C:N ratios of the 5-10 and 10-30 cm mineral soil layers (Table 4), with
293 higher N concentrations and lower C:N ratios in the deciduous stand. No significant interaction
294 effects between forest type and edge distance were observed in the mineral soil.

295 The N stocks in the mineral soil at 0-5 cm, 5-10 cm, and 10-30 cm depth were, on average, 376 (or
296 49%), 148 (or 28%), and 419 (or 25%) kg ha⁻¹ higher at the outer edge than in the forest interior,
297 respectively, the difference being significant for the 0-5 cm topsoil and marginally insignificant in
298 the other mineral soil layers (Table 3). When the deciduous and coniferous stands were considered
299 separately (Table 3), the pattern was consistent but less outspoken, with only marginally
300 insignificant differences at 0-5 cm and 5-10 cm depth in the deciduous stands and at 0-5 cm depth
301 in the coniferous stands. Summed over the total sampled mineral soil layer (0-30 cm), differences in
302 the soil N stock between outer edge and interior were larger in the deciduous forest types (1232 kg
303 ha⁻¹, i.e., 427, 176, and 629 kg ha⁻¹, for the increasing soil depth ranges) than in the coniferous ones
304 (652 kg ha⁻¹, i.e., 325, 119, and 208 kg ha⁻¹, for the increasing soil depth ranges). The N stocks in

305 the mineral soil did not differ significantly between the coniferous and the deciduous stands for any
306 of the sampling depths (Table 3).

307 Plotting the DIN leaching fluxes at 30 and 90 cm soil depth against the C:N ratio of the FH layer
308 ($C:N_{FH}$) and the mineral topsoil ($C:N_{0-5}$) (Fig. 3) showed distinct clusters related to the forest type,
309 e.g., in the deciduous stands, similar DIN leaching occurred at lower $C:N_{FH}$ ratios than in the
310 coniferous stands. In both forest types, the DIN leaching at 30 cm soil depth was not significantly
311 related to the C:N ratios. The DIN leaching at 90 cm depth in the deciduous stands was marginally
312 insignificantly related to $C:N_{FH}$ ($R^2 = 0.19$; $P = 0.097$; slope = -7.93), but significantly negatively
313 related to $C:N_{0-5}$ ($R^2 = 0.35$; $P = 0.015$; slope = -3.03), $C:N_{5-10}$ ($R^2 = 0.51$; $P = 0.002$; slope = -1.90),
314 and $C:N_{10-30}$ ($R^2 = 0.38$; $P = 0.012$; slope = -1.72) soil layers. The relationship between DIN
315 leaching at 90 cm depth and $C:N_{FH}$ disappeared when the extreme point at $C:N_{FH} = 16.4$ (see Fig. 3)
316 was removed ($R^2 = 0.02$, $P = 0.636$). In the coniferous stands, relations of DIN leaching at 90 cm
317 with neither $C:N_{FH}$ nor the C:N ratios of the mineral soil were significant. Linear regression on log-
318 transformed DIN leaching values yielded higher P and R^2 values for the relation between DIN
319 leaching at 90 cm and the C:N ratio at 0-5 cm ($R^2 = 0.48$; $P = 0.003$) and 5-10 cm ($R^2 = 0.63$; $P <$
320 0.001) of the mineral soil in the deciduous stands.

321

322 3.3. N in stems and fresh foliar material

323 The LAI and stem volume were generally higher in the first 20 m of the edges than in the interior,
324 the differences being marginally insignificant for LAI and significant for stem volume (Table 3).

325 The N concentrations in stem wood and foliar material differed not significantly between the forest
326 edge and the interior (Table 3).

327 The N stock in the foliar material (Table 3) was, on average, 3 kg N ha^{-1} (21%) higher in the first 20
328 m of the edge (mean 20 kg N ha^{-1}) than in the interior (mean 16 kg N ha^{-1}). In stem biomass, on
329 average, an additional amount of 27 kg N ha^{-1} (20%) was stored in the first 20 m of the edge

330 (average 162 kg N ha⁻¹) in comparison with the interior (average 136 kg N ha⁻¹). In the deciduous
331 stands, the mean difference was 5 kg N ha⁻¹ in foliar biomass and 22 kg N ha⁻¹ in stem biomass. The
332 differences between edge and interior in the coniferous stands amounted to 2 kg N ha⁻¹ in foliar
333 biomass and 32 kg N ha⁻¹ in stem biomass. However, the effect of edge proximity was not
334 significant, even when forest types were considered separately (Table 3). The N stock in foliar
335 material and stem wood was not significantly affected by forest type (Table 3).

336

337 **4. Discussion**

338 4.1. Forest type effect on DIN leaching

339 Higher NO₃⁻-N concentrations in soil water and DIN leaching occurred in the coniferous pine stands
340 than in the deciduous oak and birch stands, a result in line with previous studies performed in the
341 interior of similar forest types (Herrmann and others 2005; De Schrijver and others 2004, 2008) and
342 deciduous and coniferous forest types in general (von Wilpert and others 2000; Rothe and others
343 2002; De Schrijver and others 2007b). Gundersen and others (1998), MacDonald and others (2002),
344 and De Schrijver and others (2007b) demonstrated a positive relationship between DIN leaching
345 and N deposition at N saturated sites. This would imply that the forest type effect on DIN leaching
346 would be related primarily to differences in atmospheric N input. Although the species-related
347 difference in DIN leaching increases with the difference in N throughfall deposition or N load
348 (Rothe and others 2002; De Schrijver and others 2007b), stand history, N uptake, soil N retention,
349 and gaseous N losses may surpass the deposition effect (Rothe and others 2002).

350 The deciduous stands had lower C:N ratios in both the FH layer of the forest floor (C:N_{FH}) and the
351 the mineral topsoil (0-5 and 5-10 cm) than the coniferous stands, in accordance with Kristensen
352 and others (2004) and Gundersen and others (2009). Dise and others (1998), Gundersen and others
353 (1998, 2009), and MacDonald and others (2002) suggested the use of N input and the C:N ratio of
354 the forest floor or mineral topsoil to predict NO₃⁻ leaching. In our coniferous stands, DIN leaching

355 at 30 and at 90 cm were related with neither N input nor C:N ratios. However, DIN leaching at 30
356 cm related significantly with DIN throughfall deposition in the coniferous stands when only plots
357 with $C:N_{FH} \leq 25$ were considered. This result would confirm the finding of Gundersen and others
358 (1998, 2009) and MacDonald and others (2002) that in coniferous stands for organic layer $C:N \leq 25$
359 DIN leaching is related to DIN input, although the slope of DIN leaching vs. DIN throughfall is
360 higher than 1. In the deciduous stands, DIN leaching was significantly negatively related with C:N
361 ratios of the 0-5, 5-10, and 10-30 cm soil layers. Our study supports the finding of Gundersen and
362 others (2009) that a C:N value of the organic layer of 25 is not a useful threshold value to
363 discriminate between low and high risk of DIN leaching in deciduous stands. Gundersen and others
364 (2009) suggested that the C:N ratio of the mineral topsoil (0-10 cm) better reflects the N status of
365 mull type soils - in their study related to deciduous stands - and that the organic layer C:N may be
366 better suited for mor and moder type soils.

367 When analysing the effect of forest type on biogeochemical fluxes, stands should meet the
368 prerequisite of comparable climate, air pollution level, soil type, land-use history, and successional
369 stage (De Schrijver and others 2007b). Our study examined forest stands on the same soil type and,
370 within one region, under similar air quality and climate conditions. Land-use history was alike for
371 all stands down to 1770, except for the Qr1 stand, which was shortly managed as grassland around
372 1900. Soil pH, soil N, nitrification rates, and NO_3^- leaching can remain elevated and soil C
373 concentration and C:N ratios lowered in post-agricultural forests, even a century after agricultural
374 abandonment (Compton and Boone 2000; Jussy and others 2002; Gundersen and others 2009).

375 However, on sandy soils and after more than 50 years of reforestation, tree species rather than land-
376 use history affects N cycling (Compton and others 1998). In the 90-year old Qr1 stand, soil pH-KCl
377 was higher than in the other forest stands (Table 1), but C:N ratios of the mineral topsoil (on
378 average 17.7 in the 128 m plot) were comparable to those in the other oak stand Qr2 (on average
379 18.6 in the 128 m plot). Moreover, DIN leaching fluxes in the Qr1 stand were not higher than in the

380 Qr2 stand, which suggests a limited effect of the former grassland management on soil N cycling in
381 Qr1. Canopy closure was complete in all stands, but while all stands were in timber stage, the birch
382 stands Bp1 and Bp2 were in late pole stage. The trees in the birch stands are close to the
383 culmination point of height growth. Further development of these stands will be characterised by an
384 increase in N deposition input, due to increasing tree height (Draaijers 1993), and a decrease in N
385 accumulation in biomass, since after canopy closure the most N-rich biomass compartments
386 (canopy foliage) are synthesised (Richter and others 2000), and this can result in a slight increase in
387 DIN leaching (Hansen and others 2007).

388

389 4.2. Edge effect on DIN leaching

390 In comparison with the forest interior, enhanced leaching fluxes of DIN are found in the 64-m edge
391 zone of both the coniferous and deciduous forests. These differences in DIN leaching between the
392 64-m edge zone and the interior are larger than these in DIN throughfall deposition because the
393 most pronounced enhancement of DIN throughfall deposition is restricted to the first 20 m of the
394 edge while the enhancement of DIN leaching occurs between 16 and 64 m from the edge. The
395 increase in DIN soil leaching in the 64-m edge zone was greatest in the coniferous forests, and this
396 for both sampling depths, in accordance with the more pronounced edge effects on DIN throughfall
397 deposition in the coniferous forests than in the deciduous ones (Wuyts and others 2008a, b). Below
398 90 cm, an additional $14 \text{ kg N ha}^{-1} \text{ y}^{-1}$ and $21 \text{ kg N ha}^{-1} \text{ y}^{-1}$ was leached in the 64-m edge zone in the
399 deciduous and coniferous forests in comparison with the forest interior. However, for DIN leaching
400 as well as for the other variables considered (soil N and C concentrations and N stocks) we found
401 no significant difference in edge effects between the coniferous and deciduous forest types.
402 The patterns of DIN leaching and NO_3^- -N concentration did not univocally reflect the patterns of
403 DIN throughfall deposition (Fig. 2). In the *Q. robur*, *B. pendula*, and *P. nigra* stands, maximum
404 DIN throughfall deposition occurred at the front of the edges (0-2 m from the edge) behind which

405 deposition decreased with increasing edge distance (Wuyts and others 2008b; Fig. 2). In the Ps1
406 stand, edge effects on throughfall deposition of DIN were not observed because the edge of the Ps1
407 stand borders a very busy motorway: trucks passing by at < 5 m from the edge considerably disturb
408 the expected pattern of wind speed and turbulence across forest edges. We found no evidence of
409 higher NO₃⁻-N concentrations in the soil water near the exposed edges and remarkably low DIN
410 leaching occurred in the outer edge (the first 0-20 m of the edge) of the stands. At 30 cm soil depth,
411 soil DIN leaching was significantly related to the DIN throughfall deposition, but at 90 cm, the
412 relationship with throughfall was lost. The DIN leaching fluxes at 90 cm soil depth increased with
413 distance in the first 20 m of the edge until the highest values were reached at 32 and 64 m, and
414 decreased again in the forest interior at 128 m. In the first meters of an exposed beech edge,
415 Kinniburgh and Trafford (1996) reported very low NO₃⁻-N concentrations in the pore water beneath
416 the rooting zone; however, concentrations were increased at 30-100 m from the edge, after which
417 they decreased again. These results and those of Spangenberg and Kölling (2004) agree with our
418 findings on leaching of DIN. The non-uniform response of DIN leaching to edge proximity is
419 probably the result of a complex interaction of several processes: (i) N uptake by trees and the
420 understory vegetation, (ii) gaseous N loss, (iii) N immobilisation in the soil, and (iv) dissolved
421 organic N (DON) leaching. However, because not all possible N pathways were actually measured
422 in this study, we stress the importance of further research to verify our assumptions.

423 We avoided the influence of N uptake by understory vegetation on DIN leaching, except in the Pn2
424 stand, where lower DIN leaching at the edges may be related to N uptake by grasses at the edge.

425 Also, in stand Qr1, the brambles that remained outside the cleared zone around the lysimeters might
426 still have affected DIN leaching in the first 20 m of the edge to a minor extent. Yet, the effect of
427 these brambles can be assumed negligible as their abundance was low. The stand volume and the
428 LAI were higher in the first 20 m of the edge, but differences in N stock in these biomass
429 compartments were not significant. So, though we did not measure N stocks in all biomass

430 compartments, these results do not confirm the hypothesis by Spangenberg and Kölling (2004) on N
431 sequestration in trees and understory biomass and instead indicate that the role of higher N
432 accumulation in biomass in lower DIN leaching in the outer edge is limited.

433 Production of NO and N₂O, primarily driven by microbial processes, depends on soil temperature,
434 moisture, pH, and N input (Butterbach-Bahl and others 2002; Schindlbacher and others 2004;
435 Pilegaard and others 2006). From edge gradients in N throughfall deposition, pH, and microclimate
436 (e.g., soil temperature; Chen and others 1995), elevated N loss through NO and N₂O emission can
437 be expected at edges.

438 N concentrations in the mineral topsoil (0-5 cm) were significantly enhanced at the edges and the N
439 stock in the total sampled mineral soil (0-30 cm) was, on average, 951 kg N ha⁻¹ higher in the first
440 20 m of the edge than in the interior. Processes that are believed to be involved in N retention in the
441 soil are (a)biotic fixation of NO₃⁻ and NH₄⁺ in or on soil organic matter (Magill and others 1997;
442 Aber and others 1998; Davidson and others 2003; Micks and others 2004; Morier and others 2008).
443 NH₄⁺ adsorption in the mineral soil is positively related to soil pH and exchangeable amounts of
444 Mg²⁺ (Kothawala and Moore 2009), which were both enhanced in the mineral topsoil (0-5 cm) in
445 the first 32 m of our edges (data on Mg²⁺ in Fig. 4). Moreover, Hayes (2002) suggested that at forest
446 edges a complex interaction of processes is involved in soil N retention, such as increased litter
447 decomposition and higher biomass input through litterfall. The magnitude and extent of the edge
448 effects on these processes may differ between edges with dissimilar structure, species composition,
449 exposition, and type of adjacent land cover (Hayes 2002; Malmivaara-Lämsä and others 2008). We
450 presume that both litter decomposition and litterfall input are increased at the edges, but that the
451 edge effect on the decomposition rate penetrates the forest deeper than the effect on litterfall. Litter
452 decomposition depends, among other factors, on soil biological activity. Higher soil temperature,
453 occurring up to 50-60 m from the edge (Chen and others 1995; Heithecker and Halpern 2007), and
454 higher Ca content in the litter (Fig. 4), as a result of higher base cation deposition (Wuyts and others

455 2008a), could favour the abundance and activity of soil and litter organisms and increase
456 decomposition rates at the edges (Reich and others 2005; Hobbie and others 2007). Hence, we
457 hypothesize that with an increase in litter decomposition, less N is retained in the forest floor and
458 more N is flushed to the mineral soil up to 64 m from the edge. At our edges, though LAI was
459 higher, the FH layer dry weight at the edge was on average 28% lower than in the interior, which
460 can be related to higher decomposition rates but also to blow up of litter by wind from the edge to
461 the interior. Next, we presume the presence of an edge effect on litterfall input that penetrates the
462 edge less far than the edge effect on decomposition. In the outer 20 m of the edges, the biomass of
463 fresh leaves (LAI), and presumably also the litterfall flux to the forest floor, was higher than further
464 down the edge and in the interior. The soil organic matter content was enhanced near the edges,
465 indicating a higher potential for N immobilisation in or on soil organic matter. Also Marchand and
466 Houle (2006) observed a higher soil organic matter content in the first 20 m of a forest edge. In the
467 outer edge of a *P. nigra* stand near Pn2, Vandenbruwane (2008) found higher DON and dissolved
468 organic C (DOC) fluxes from the forest floor to the mineral soil and higher retention of DON and
469 DOC in the upper 70 cm of the mineral soil in the outer edge than in the interior (13 vs. 7 kg N ha⁻¹
470 y⁻¹). Next to higher DON retention, Vandenbruwane (2008) observed higher DON leaching fluxes
471 below the main rooting zone in the outer edge than in the interior (5 vs. 3 kg N ha⁻¹ y⁻¹). Hence, the
472 proportion of organic N in the total N leaving the system via leaching is increased at the outer edge
473 in comparison with the interior.

474 We can conclude that the C:N ratio and the DIN input with throughfall deposition cannot serve as
475 indicators or predictors for DIN leaching in forest edges: at the outer edges, higher DIN throughfall
476 deposition and lower C:N ratios were observed, yet lower DIN leaching was detected in comparison
477 with the plots at 32 and 64 m from the edge, particularly at 90 cm soil depth. Although we found a
478 significant relationship between DIN throughfall and DIN leaching at 30 cm in the deciduous stands
479 (and in the coniferous stands for $C:N_{FH} \leq 25$), such relationships were no longer appreciable at 90

480 cm soil depth. Between DIN leaching and C:N ratio, only significant relationships were found in the
481 deciduous stands, and the variability in DIN leaching explained by the C:N ratio of the mineral soil
482 was low.

483

484 4.3. Methodological considerations

485 The use of only one suction cup lysimeter per depth and distance plot may not take into account the
486 spatial variability in DIN leaching at a given edge distance. Kohlpaintner and others (2009) found
487 that only understory vegetation coverage, distance to the nearest tree, and tree basal area within a
488 radius of 4 m contributed to explain the variability in NO_3^- leaching in a mature Norway spruce
489 stand. Given that in our stands, (i) understory vegetation was absent or removed more than half a
490 year before sampling (in exception of the Pn2 stand, where a homogeneous grass cover occurred at
491 the edge), and (ii) forest stands were homogeneous in basal area, with exception of the edge front,
492 we expect spatial variability in DIN leaching within one distance plot to be limited. The good
493 agreement (high correlation) in NO_3^- -N concentrations and DIN leaching between the lysimeters of
494 the two sampling depths (with at least 1 m intermediate distance) also suggests a relatively low
495 spatial variability.

496 The DIN leaching fluxes were calculated with water percolation fluxes estimated from the chloride
497 mass balance. Although Cl^- acts as a conservative element in the soil of undisturbed forests in the
498 long term (Kauffman and others 2003) and is even close to conservative within 2 years (Rosenqvist
499 and others 2010), imbalances between Cl^- input and output within one year and periods of sea salt
500 peaks may occur, causing the water percolation flux and DIN leaching to be under- or
501 overestimated. In this study, the water percolation fluxes might be overestimated due to a peak in
502 sea salt deposition that occurred during a storm event on 25 November 2005, particularly in the
503 region the closest to the sea (region 1). Although this peak in Cl^- deposition occurred in the
504 beginning of the year of sampling, it may not have passed completely through the soil by the end of

505 the sampling period. However, the similar trend in the measured NO_3^- -N concentrations in the soil
506 solution as in the derived DIN leaching fluxes and the large number of studied stands and distance
507 plots reduce the uncertainty in the edge pattern of DIN leaching related to the use of the chloride
508 mass balance.

509

510 **5. Conclusions**

511 In 64 m edge zones of deciduous and coniferous forests, DIN leaching fluxes were higher than in
512 forest interiors. DIN leaching was higher and exhibited the largest edge effects in coniferous forests.
513 DIN leaching at 90 cm soil depth did not relate with DIN throughfall deposition; in the first 20 m
514 (i.e., the outer edge), DIN leaching was reduced in comparison with further down the edge (20-64
515 m). Increased soil N retention is probably one of the processes involved in this local decrease in
516 DIN leaching. Our results point towards a complex of edge effects on biogeochemical processes, a
517 result of interactions between edge effects on N deposition and on base cation deposition,
518 microclimate, and tree dynamics (e.g., growth and litterfall). The course and extent of these edge
519 effects may differ according to structure, exposition, species composition, and adjacent land cover,
520 complicating the generalisation of edge patterns in biogeochemical and ecological studies.

521

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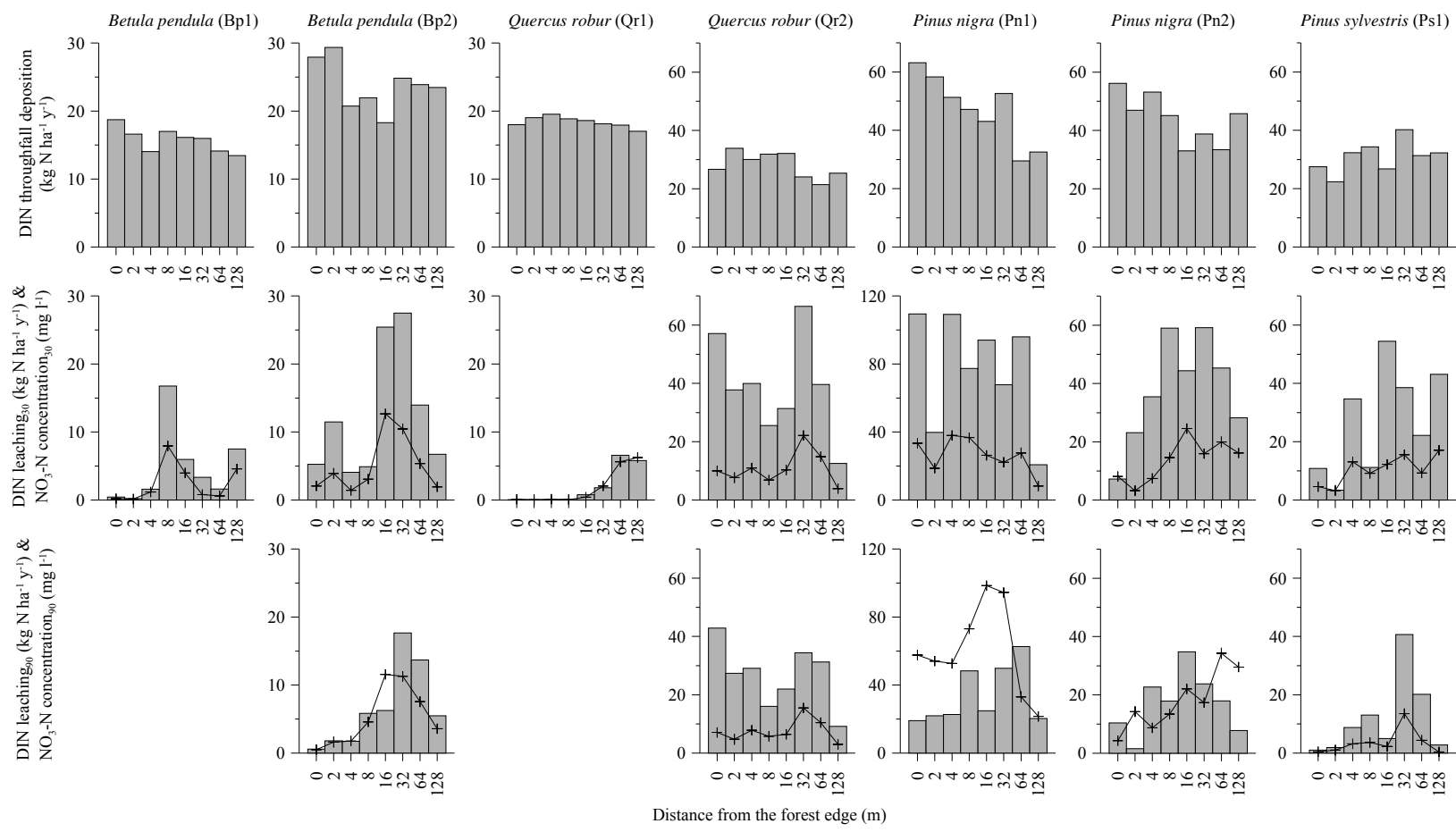
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689 **Appendix**

690 Throughfall deposition and soil leaching fluxes at 30 cm (DIN leaching₃₀; indicated by bars) and 90 cm (DIN leaching₉₀; bars) soil depth of DIN ($\text{NO}_3^- + \text{NH}_4^+$; $\text{kg N ha}^{-1} \text{y}^{-1}$)
 691 and volume-weighted mean NO_3^- -N concentrations (mg l^{-1} ; crosses) in the soil solution along transects across the studied forest edges. Soil leaching fluxes at 90 cm were not
 692 measured in the stands Qr1 and Bp1 due to high ground water tables. Data on throughfall deposition were previously published by Wuyts and others (2008b), except for Ps1.
 693 Note the higher y-axis limit for the stands Qr2, Pn1, Pn2 and Ps1.



694

695 **Figure legends**

696 Fig. 1. Maps indicating location of Flanders within Europe and the eight studied forest stands in the two
697 regions (see Table 1 for explanation of site codes).

698

699 Fig. 2. Mean throughfall deposition of DIN ($\text{NH}_4^+ + \text{NO}_3^-$; $\text{kg N ha}^{-1} \text{ y}^{-1}$) and DIN soil leaching fluxes (kg N
700 $\text{ha}^{-1} \text{ y}^{-1}$) and NO_3^- -N concentrations (mg l^{-1}) in the soil solution at 30 (DIN leaching₃₀ and NO_3^- -N
701 concentration₃₀) and 90 cm (DIN leaching₉₀ and NO_3^- -N concentration₉₀) soil depth along transects across the
702 deciduous and coniferous forest sites. Error bars indicate standard errors. Note the altered y-axis limits.

703

704 Fig. 3. DIN leaching ($\text{kg N ha}^{-1} \text{ y}^{-1}$) at 30 and 90 cm soil depth plotted against the throughfall deposition of
705 DIN ($\text{kg N ha}^{-1} \text{ y}^{-1}$) and the C:N ratio of the fermentation and humus layer of the forest floor (C:N_{FH}) and of
706 the 0-5 cm mineral topsoil ($\text{C:N}_{0.5}$) in the separate distance plots in the deciduous stands (●) and the
707 coniferous ones (+). Indicated are the (marginally) significant relationships. All data included, except for the
708 Ps2 stand.

709

710 Fig. 4. Mean N (g kg^{-1}) and C (%; determined by loss on ignition) concentrations and C:N ratios of the
711 fermentation and humus layer (FH) of the forest floor and the 0-5 cm, 5-10 cm, and 10-30 cm of the mineral
712 soil and mean Ca and Mg concentrations of the FH layer ($\text{mol}_c \text{ kg}^{-1}$) and exchangeable concentrations in the
713 top 5 cm of the mineral topsoil ($\text{mmol}_c \text{ kg}^{-1}$) along transects across the forest edges of all the studied forest
714 stands. Error bars indicate standard errors. Note the altered y-axis limits for the forest floor and the different
715 soil depths.

716 **Tables**

717 Table 1.

Site code	Location		Tree species	Age (y)	SN (ha ⁻¹)	BA (m ² ha ⁻¹)	V (m ³ ha ⁻¹)	H (m)	H _{dom} (m)	LAI		pH-KCl
										Summer	Winter	
Bp1	51°09'22"N	03°04'48"E	<i>Betula pendula</i>	30-40	3628	26	194	11.2	18.2	1.36	0.61	2.93
			<i>Sorbus aucuparia</i> ⁺		389	2	5					
Bp2	51°25'56"N	05°00'31"E	<i>Betula pendula</i>	20-30	2715	13	74	8.1	13.8	1.33	0.31	3.03
			<i>Quercus robur</i> ^A		111	3	34					
Qr1	50°52'08" N	03°27'59"E	<i>Quercus robur</i>	90	187	31	343	24.2	26.4	1.90	0.79	3.35
Qr2	51°24'44" N	05°02'45"E	<i>Quercus robur</i>	68	135	23	221	21.1	22.4	1.88	0.40	2.88
			<i>Pinus nigra</i> ssp.									
Pn1	51°08'26"N	03°06'36"E	<i>laricio</i>	65	388	36	336	19.2	22.0			2.91
Pn2	51°26'37"N	05°05'14"E	<i>Pinus nigra</i> ssp. <i>nigra</i>	43	1162	55	488	17.3	19.1	1.80	-	2.89
Ps1	51°10'11"N	03°09'36"E	<i>Pinus sylvestris</i>	80	458	29	287	19.8	23.2	1.82	-	2.95
Ps2	51°24'45"N	05°02'39"E	<i>Pinus sylvestris</i>	76	195	20	183	19.4	21.1	1.03	-	2.86

718 ⁺: sum of basal area of other species (*Quercus robur* and *Rhamnus frangula*) < 1 m² ha⁻¹719 ^A: sum of basal area of other species (*Robinia pseudo-acacia* and *Quercus rubra*) < 1 m² ha⁻¹

720 Table 2.

	TF	Leaching 30 cm		Leaching 90 cm	
	P value	[NO ₃ ⁻ -N] P value	DIN flux P value	[NO ₃ ⁻ -N] P value	DIN flux P value
RMA					
Factor					
Distance to the forest edge (8)	0.021	0.161	0.281	0.500	0.341
Forest type (2)	0.015	0.054	0.041	0.413	0.462
Distance x forest type	0.211	0.688	0.379	0.960	0.624
Edge proximity (2)		0.339	0.090	0.391	0.003
Forest type (2)		0.002	0.005	0.328	0.383
Edge proximity x forest type		0.788	0.440	0.391	0.223
WSR					
	TF (kg N ha ⁻¹ y ⁻¹)	[NO ₃ ⁻ -N] (mg l ⁻¹)	DIN flux (kg N ha ⁻¹ y ⁻¹)	[NO ₃ ⁻ -N] (mg l ⁻¹)	DIN flux (kg N ha ⁻¹ y ⁻¹)
Forest type					
Deciduous	21 ± 2	7.0 ± 3.2	18 ± 10	9.8 ± 1.0	21 ± 9
0-64 m	20 ± 3	4.2 ± 0.9	8 ± 2	3.3 ± 0.3	7 ± 2
128 m					
Coniferous					
0-64 m	39 ± 3	19 ± 4	55 ± 15	34 ± 20	31 ± 8
128 m	37 ± 4	14 ± 3	31 ± 7	17 ± 9	10 ± 5

721 Table 3.

RMA	FH layer	Mineral soil - N stock			Stem wood			Foliar material		
	dry weight	0-5 cm	5-10 cm	10-30 cm	Biomass	[N]	N stock	LAI	[N]	N stock
	P value	P value	P value	P value	P value	P value	P value	P value	P value	P value
Factor										
Edge proximity (2)	0.044	0.003	0.078	0.091	0.043	0.352	0.131	0.058	0.394	0.188
Forest type (2)	0.711	0.542	0.232	0.351	0.117	0.003	0.830	0.808	0.009	0.943
Edge prox. x forest type	0.431	0.530	0.692	0.844	0.294	0.049	0.729	0.880	0.010	0.454
WSR	FH layer	Mineral soil - N stock (kg N ha ⁻¹)			Stem wood			Foliar material		
	dry weight (kg m ⁻²)	0-5 cm	5-10 cm	10-30 cm	Biomass (m ³ ha ⁻¹)	[N] (mg kg ⁻¹)	N stock (kg N ha ⁻¹)	LAI (m ² m ⁻²)	[N] (g kg ⁻¹)	N stock (kg N ha ⁻¹)
Forest type										
Deciduous										
0-20 m	7.5 ± 1.1	1218 ± 159	778 ± 118	2259 ± 225	193 ± 54	1140 ± 67	158 ± 42	2.0 ± 0.1	26.5 ± 0.5	20.2 ± 1.4
128 m	9.8 ± 0.4	791 ± 142	602 ± 85	1630 ± 346	176 ± 55	987 ± 29	126 ± 40	1.6 ± 0.2	25.2 ± 0.9	15.1 ± 1.7
Coniferous										
0-20 m	7.5 ± 0.9	1056 ± 130	584 ± 74	1980 ± 275	370 ± 81	629 ± 63	166 ± 47	1.9 ± 0.4	15.6 ± 1.7 [†]	18.8 ± 4.6
128 m	11.6 ± 2.0	732 ± 90	466 ± 118	1772 ± 252	327 ± 62	693 ± 91	145 ± 46	1.5 ± 0.2	17.5 ± 1.4 [†]	17.2 ± 4.1

722 †: current-year needles: 17.9 ± 1.9 and 19.8 ± 1.6 g N kg⁻¹ at 0-20 m and 128 m, resp.

723 Table 4.

Variable	Factor	FH	Mineral soil		
			0-5 cm	5-10 cm	10-30 cm
N	Distance to the forest edge (8)	0.011	<0.001	0.130	0.078
	Forest type (2)	0.726	0.037	0.030	0.441
	Distance x forest type	0.631	0.205	0.086	0.811
C (LOI)	Distance to the forest edge (8)	0.008	<0.001	0.021	0.014
	Forest type (2)	0.044	0.517	0.223	0.570
	Distance x forest type	0.578	0.312	0.086	0.313
C:N	Distance to the forest edge (8)	0.002	0.023	0.516	0.125
	Forest type (2)	<0.001	0.003	0.080	0.062
	Distance x forest type	0.313	0.171	0.163	0.933

724 **Table legends**

725 Table 1. Studied sites (forest stands) with indication of their location, the tree species, and the mean leaf area
726 index (LAI) and pH-KCl (1 M KCl) of the mineral topsoil (0-5 cm) in the forest interior. The mean age, stem
727 number (SN), mean and dominant tree height (H and H_{dom}, respectively), basal area (BA), and stand volume
728 (V) are given for the dominant tree species of each stand.

729

730 Table 2. Overview of (i) the repeated-measures ANOVA outcome (RMA; P values) and (ii) the weighted-
731 mean values (\pm standard error) in the first 64 m of the forest edge (0-64 m) and in the forest interior (128 m)
732 in the deciduous and coniferous forest stands for the NO₃⁻-N concentrations in soil water ([NO₃⁻-N]) and
733 DIN leaching at 30 and 90 cm soil depth. *Italic*: marginally insignificant difference (0.05 < P < 0.10)
734 according to Wilcoxon signed ranks (WSR) tests.

735

736 Table 3. Overview of (i) the repeated-measures ANOVA outcome (RMA; P values) and (ii) the weighted-
737 mean values (\pm standard error) in the outer forest edge (0-20 m) and in the forest interior (128 m) in the
738 deciduous and coniferous forest stands for the dry weight of the fermentation and humus layer (FH) of the
739 forest floor, the N stock in the mineral soil (depths 0-5 cm, 5-10 cm, and 10-30 cm), the biomass, N
740 concentration ([N]), and N stock in tree stems, and the leaf area index (LAI), N concentration, and N stock in
741 tree foliar material (leaves and old needles). *Italic*: marginally insignificant (0.05 < P < 0.10) difference
742 between edge and interior according to Wilcoxon signed ranks (WSR) tests.

743

744 Table 4. Significance (P value) of edge effects (factor 'distance to the forest edge'), forest type effects, and
745 their interaction ('Distance x forest type') on the C (estimated by loss on ignition, LOI) and N concentration
746 and the C:N ratio of the fermentation and humus layer (FH) of the forest floor and the mineral soil according
747 to a repeated-measures ANOVA (**bold**: P < 0.05).

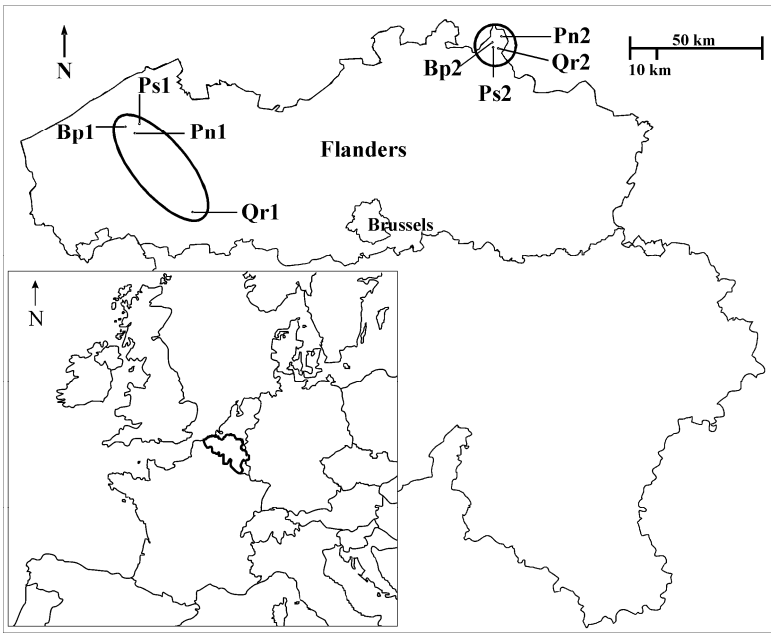


Fig. 1

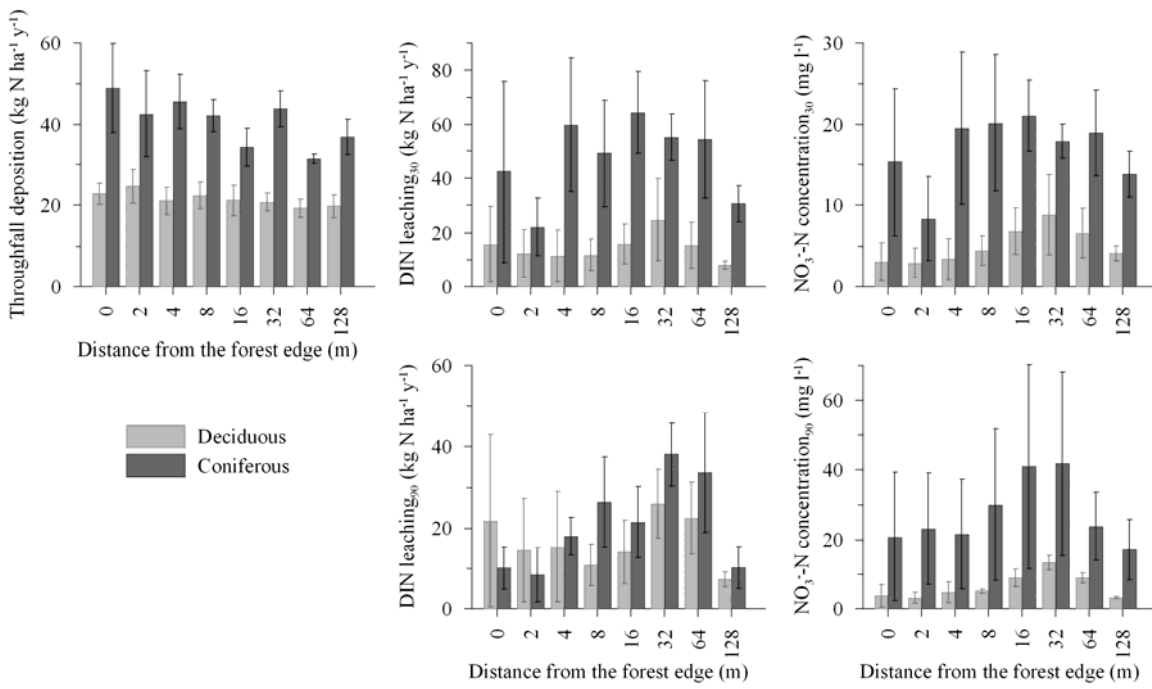


Fig. 2

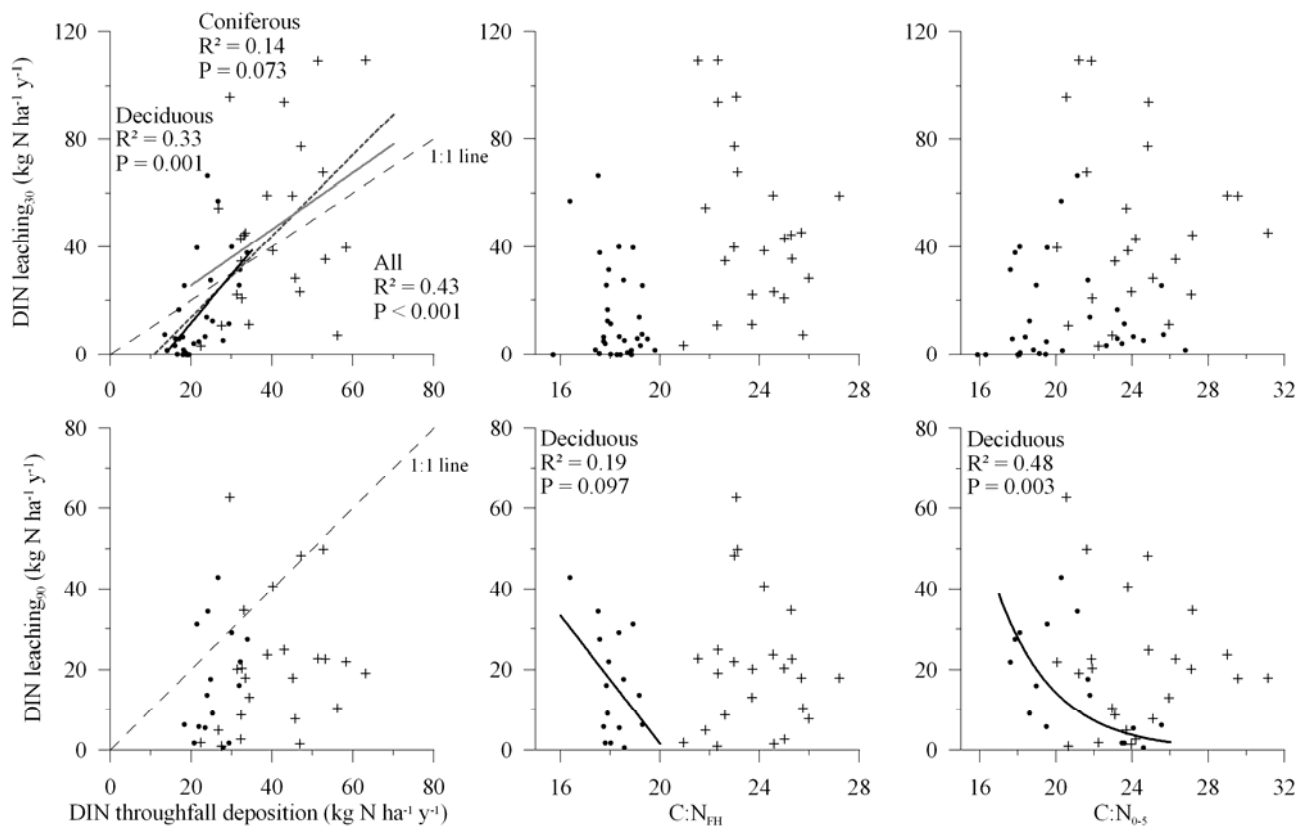


Fig. 3

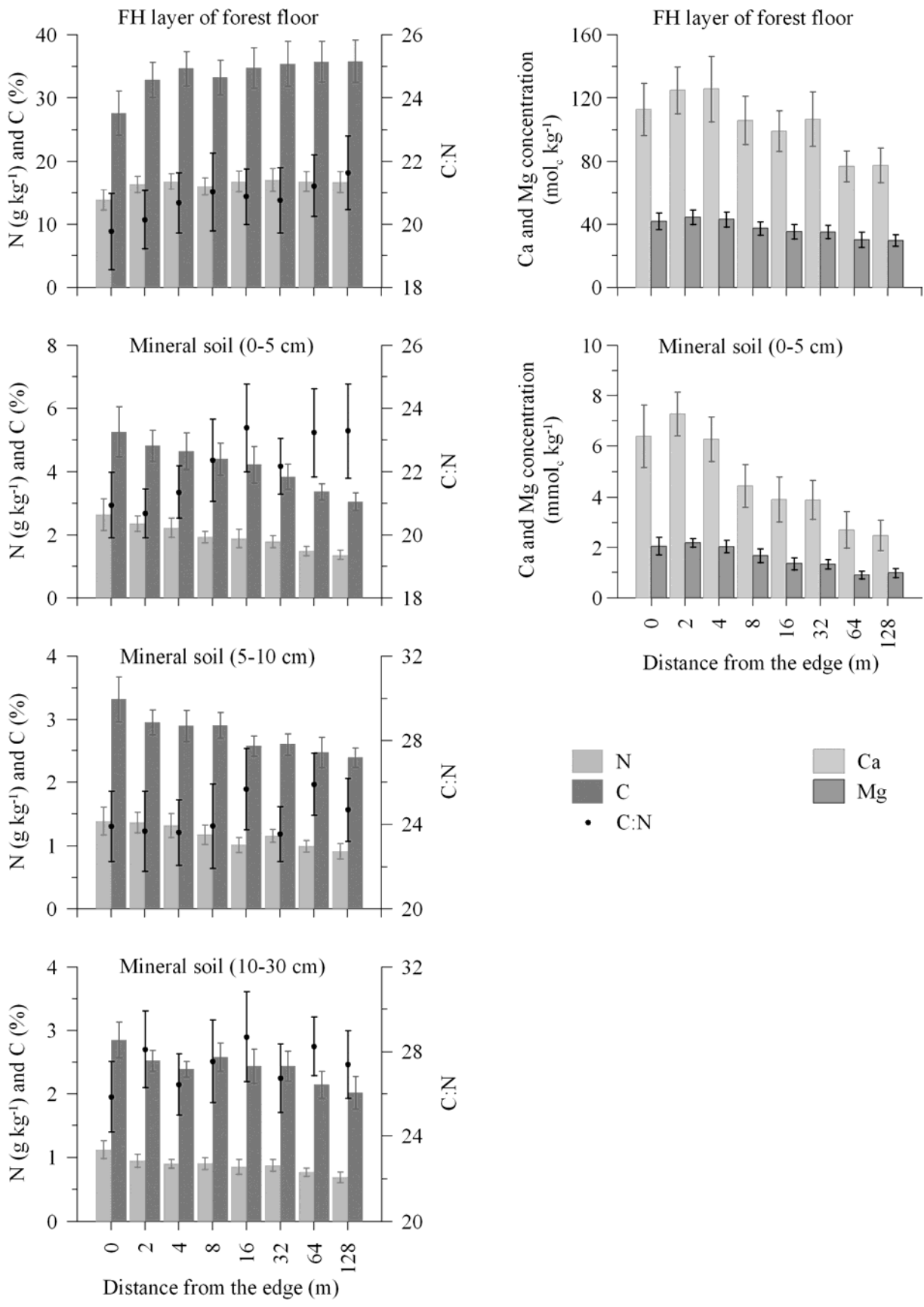


Fig.4