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Peatland records of solar activity

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Rainwater-fed Holocene raised bog deposits in temperate climate zones are valuable archives of solar activity fluctuations and related climate changes.

Past hydrological changes can be derived from the species composition and the degree of decomposition of peat-forming plants in raised bog deposits. More than 120 years ago Axel Blytt described transitions in raised bog deposits between dark, highly decomposed peat and light colored peat as evidence for climate change. Nowadays, the radiocarbon (¹⁴C) method is generally used for dating Holocene climate-induced shifts in raised bog deposits. Yet radiocarbon offers more than the possibility to precisely date organic material. The changing atmospheric radiocarbon levels in the past are also a proxy for solar activity. We have found strong links between large ¹⁴C fluctuations and major climate shifts.

¹⁴C wiggle-match dating and the Sun

The radioactive carbon isotope ¹⁴C is produced under the influence of cosmic rays. When formed, the cosmogenic isotope radiocarbon is oxidized to ¹⁴CO₂, which becomes part of the carbon cycle. When plants and animals die, the uptake of carbon ceases and the decay of radiocarbon (half-life 5730 ± 40 years) to deduce age. Radiocarbon "years" are different from calendar years because the production of ¹⁴C in the past was not constant, mainly due to changes in solar activity and related cosmic ray intensity. The ¹⁴C timescale is calibrated by measuring the ¹⁴C content of tree rings, dated precisely by means of dendrochronology. Changes in solar activity underlie fluctuations ("wiggles") in

the ¹⁴C calibration curve. Calibration of a single radiocarbon date usually yields an irregular probability distribution in calendar age, quite often over a long time interval. This is problematic in paleoclimatological studies, especially when a precise temporal comparison between different climate proxies is required. However, closely spaced sequences of (uncalibrated) ¹⁴C dates of peat deposits also display the wiggles, which can be fitted to the wiggles in the radiocarbon calibration curve.

The practice of dating peat samples using ¹⁴C "wiggle-match dating" has greatly improved the precision of radiocarbon chronologies (van Geel and Mook, 1989; Kilian et al., 1995; Blaauw et al., 2003). By wiggle-matching ¹⁴C measurements, high-precision calendar-age chronologies for peat sequences can be generated, and several studies indicate that peatland surface wetness often increased together with rapid increases in atmospheric production of ¹⁴C during the early Holocene, the Subboreal-Subatlantic transition (at ca. 2.8 cal ka BP; Fig. 1 and 2), and the Little Ice Age. Peat records from The Netherlands, the Czech Republic, the UK and Denmark show this phenomenon (Mauquoy et al., 2002a, b; Speranza et al., 2002; van der Plicht et al., 2004; van Geel et al., 1996, 1998). Because the production of radiocarbon is regulated by solar activity, periods of increased peatland surface wetness have been interpreted as evidence for solar forcing of climate change

(effects of sudden declines of solar activity).

Subboreal/Subatlantic transition and Little Ice Age

One of the most important climate shifts during the Holocene is the Subboreal/Subatlantic transition, which occurred at ca. 2.8 cal ka BP (Fig. 1 - 3). It had strong socio-economic impacts in areas that were marginal from a hydrologic point of view. In lowland regions in The Netherlands, the climate shifted to cooler and wetter conditions, causing a rapid and substantial rise of the groundwater table, transforming arable land into wetland, where peat growth started. Farming communities living in such areas were forced to migrate because they could no longer produce sufficient food (van Geel et al., 1996).

In Figure 3, the landscape development in the northern Netherlands is shown for three successive phases. At the start of the second phase, coincident with an abrupt decline of solar activity, the atmospheric circulation changed, leading to cooler and wetter climate conditions. The rise of the water table forced farmers to migrate to well-drained areas in the northern Netherlands where salt marshes offered them new fertile land.

In south-central Siberia, archeological evidence suggests an acceleration of cultural development and a sudden increase in density and geographic distribution of the nomadic Scythian population after 2.8 cal ka BP. Van Geel et al. (2004) hypothesized a relationship with an abrupt climatic shift towards increased humidity (equatorward relocation of mid-latitude storm tracks). The hypothesis is supported by palynological evidence. Areas that were initially hostile semi-deserts were converted into steppe landscapes with a high biomass production and carrying capacity. Newly available steppe areas could be utilized by herbivores, making them attractive for nomadic tribes. The equestrian Scythian culture expanded from Central Asia, and increased population density was a stimulus for westward migration towards south-eastern Europe.

Evidence for changes in peatland surface wetness, which coincide with grand solar minima during the Little Ice Age (Wolf, Spörer, Maunder and Dalton Mini-



Figure 1: The Subboreal-Subatlantic transition in a raised peat profile in The Netherlands.

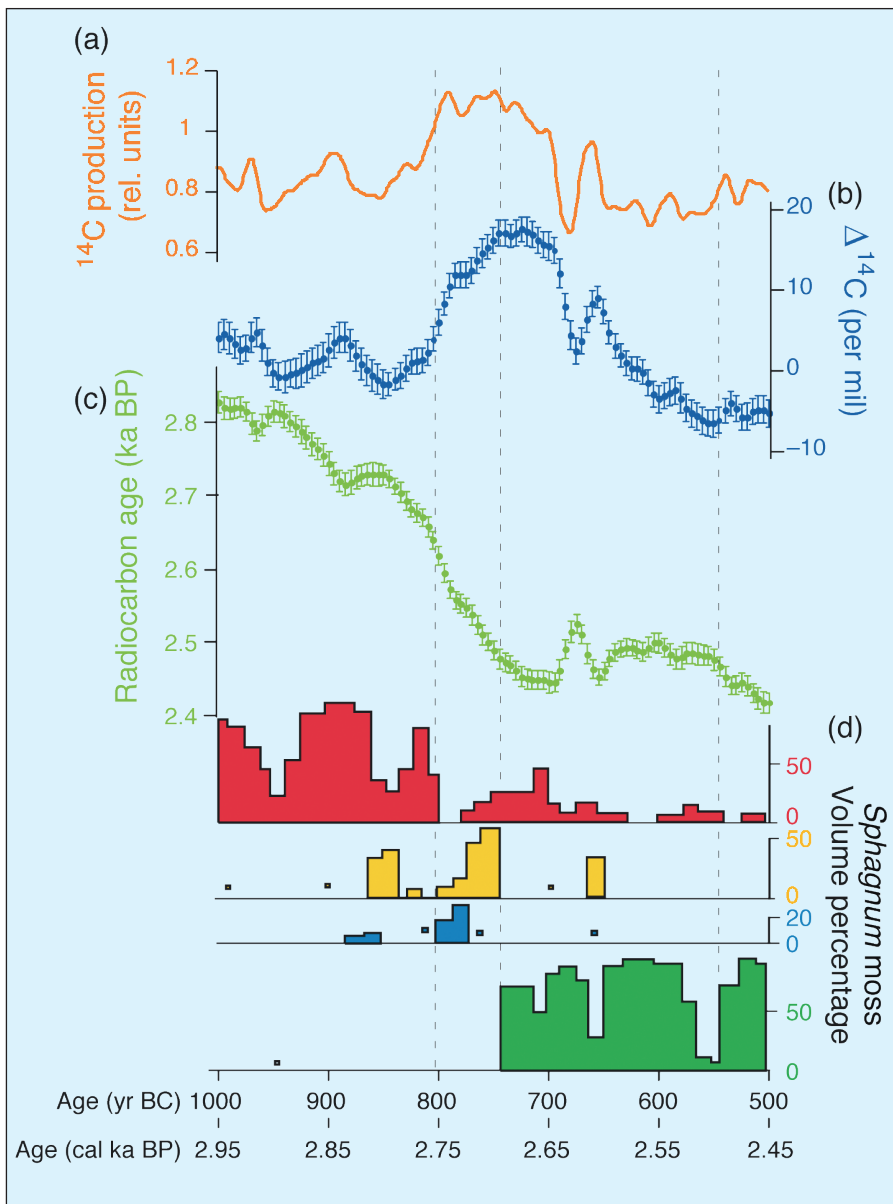


Figure 2: **a)** Inferred changes in ^{14}C production (yellow). **b)** $\Delta^{14}\text{C}$ is the deviation of the atmospheric $^{14}\text{C}/^{12}\text{C}$ ratio from a standard value in per mil (Stuiver and Pollach, 1977). The interval between ca. 2.8 and 2.7 cal ka BP represents the phase of extreme climate at the start of the Subatlantic period: the rise of $\Delta^{14}\text{C}$ and the ^{14}C production curve are related to a temporary decline of solar activity; **c)** The ^{14}C calibration curve (Reimer et al., 2004). **d)** Changes in composition of Sphagnum moss species during the Subboreal-Subatlantic transition (Engbertsdijsveen, eastern Netherlands; after van Geel et al., 1996; Beer and van Geel, 2008) reflect a sharp transition from relatively warm-dry (*S. section acutifolia* (red)) to cool wet conditions (*S. cuspidatum* (blue), *S. papillosum* (yellow), *S. imbricatum* (green)). Figure modified from Beer and van Geel, 2008).

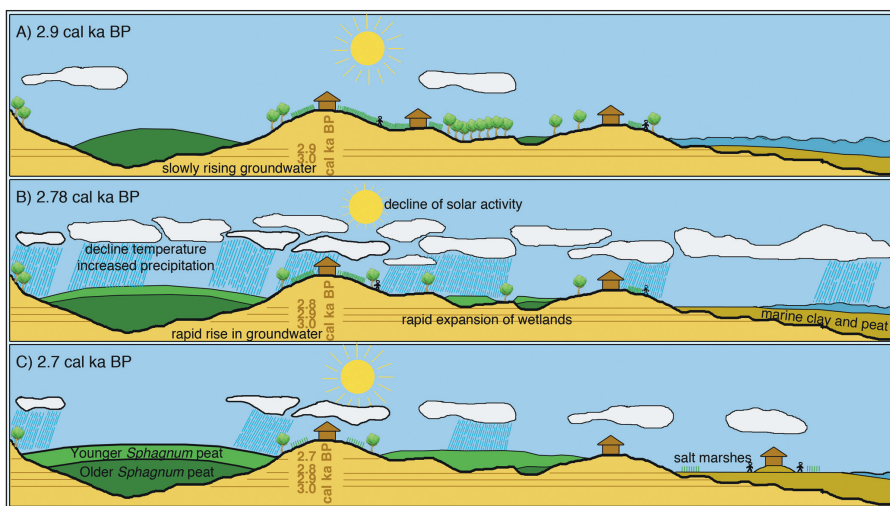


Figure 3: Landscape effects of solar forcing of climate change just before, during and directly after the Subboreal-Subatlantic transition in the northern Netherlands (modified from van Geel et al., 1998; Beer and van Geel, 2008).

ma), is presented in Mauquoy et al. (2002a, b) and Mauquoy et al. (2008). Replicate cores were investigated from one site in Denmark (2 peat profiles from Lille Vildmose) and northern England (3 replicate cores from Walton Moss). Paleohydrological reconstructions based on plant macrofossil and testate amoebae at a second site in northern England also record Little Ice Age climatic cooling (Mauquoy et al., 2008).

These increases in peatland surface wetness also coincide with some changes in solar activity; they are coeval with increases in the production of ^{14}C . All of the peat profiles possess precise ^{14}C wiggle-match dated chronologies, and the chronology from Lille Vildmose offers a very high-resolution record of climate change spanning the last 1.5 ka, with peat accumulation rates as high as 2–3 yr cm^{-1} . The study by Mauquoy et al. (2008; see also Chambers et al., this issue) indicates a short lag (10–40 years) between the start of the rise in ^{14}C production and the hydrological response of the peat bog. We do not know the reason for this. Analyses of further peat profiles are needed to replicate this finding. Additional evidence for solar forcing of climate change comes from studies of marine sediments, stalagmites, ice cores and lake sediment, showing the sensitivity of the climate system to relatively small changes in solar activity (Bond et al., 2001; Neff et al., 2001; Magny, 2004, 2007).

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B. van Geel and D. Mauquoy

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