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Jean-François Ghienne, Abderamane Moussa, Abakar Saad,
Barnabé Djatibeye and Hissein Mahamat Youssouf

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Volume 355 (2023), p. 63-84

Published online: 16 January 2023

<https://doi.org/10.5802/crgeos.180>



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www.centre-mersenne.org

e-ISSN : 1778-7025



Original Article — Petrology, Sedimentology

The Ordovician strata of the Ennedi Plateau, northeastern Chad (Erdi Basin)

*Les grès ordoviciens du plateau de l'Ennedi, NE du Tchad
(Bassin des Erdis)*

Jean-François Ghienne^{®*, a}, Abderamane Moussa^b, Abakar Saad^b, Barnabé Djatibeye^b
and Hissein Mahamat Youssouf^c

^a Institut Terre et Environnement de Strasbourg, UMR7063, CNRS – Université de Strasbourg, 5, rue R. Descartes, 67000 Strasbourg, France

^b Laboratoire Hydro-Géosciences and Réservoirs, Département de Géologie et Paléontologie, Université de N'Djamena, BP 1027, N'Djamena, Chad

^c Ministère du Pétrole du Chad, N'Djamena, Chad

E-mails: ghienne@unistra.fr (J.-F. Ghienne), mellah03@yahoo.fr (A. Moussa), saabakar8@gmail.com (A. Saad), barnabdjatibeye24@yahoo.com (B. Djatibeye), youssoufhissein@gmail.com (H. Youssouf)

Abstract. In the Ennedi Plateau, northeastern Chad, Cambrian(?)–Ordovician to Carboniferous strata represent one of the most proximal succession of the north-facing Gondwana platform. This study presents preliminary results related to the *c.* 200 m-thick Lower Sandstones and overlying Bedo Formation, which essentially record Early to Middle Ordovician flooding events and the end-Ordovician glaciation and subsequent deglaciation. In the Lower Sandstones, fluvial-dominated deposits (prevailing cross-stratification, NW to NNE paleocurrent trends) include intervening high-frequency marine incursions (tidal deposits, ichnofacies including *Cruziana* and *Arthropycus*). Above an erosional unconformity, the end-Ordovician glacial record consists of fluvioglacial conglomeratic deposits atop the Lower Sandstones, which are in turn superimposed by the fine-grained Bedo Formation. The latter interval, onlapping an exhumed erosion surface showing north-trending km-scale glacial lineations, shows glaciomarine deposits only at its base and possibly include Lower Silurian deposits in its upper part.

Résumé. Dans le massif de l'Ennedi, au NE du Tchad, les dépôts s'étageant du Cambrien(?)–Ordovicien au Carbonifère représentent l'une des successions les plus proximales de la plate-forme « nord-gondwanienne ». Cette étude présente quelques résultats préliminaires sur les Grès Inférieurs, épais d'environ 200 m, et la Formation de Bedo sus-jacente. Cet intervalle stratigraphique enregistre essentiellement les événements transgressifs de l'Ordovicien inférieur à moyen, puis la glaciation, ainsi que la déglaciation consécutive, de la fin de l'Ordovicien. Dans les Grès Inférieurs, les dépôts à dominante fluviatile (dominance des faciès à laminations obliques, orientation des courants dans

* Corresponding author.

un secteur NW à NNE) comprennent des incursions marines à haute fréquence (dépôts de marée, ichnofaciès à *Cruziana* et *Arthropycus*). Au-dessus d'une discordance d'érosion venant tronquer les Grès Inférieurs, la succession glaciaire débute avec des dépôts conglomératiques fluvioglaciaires et termine avec une unité à dominante argileuse (Formation de Bedo). Cette dernière, qui scelle une surface d'érosion exhumée à linéations glaciaires orientées vers le nord et d'échelle kilométrique, montre des dépôts glaciomarins uniquement à sa base. Elle pourrait inclure dans sa partie sommitale des dépôts du Silurien inférieur.

Keywords. Erdi Basin, Palaeozoic, Africa, Glaciation, Sandstone.

Mots-clés. Paléozoïque, Plate-forme, Grès, Afrique, Glaciation.

Manuscript received 17 May 2022, revised 3 October 2022, accepted 2 November 2022.

1. Introduction

The Ennedi Plateau is a sandstone plateau in north-eastern Chad (Figure 1), which exposes a Palaeozoic succession expanding from the Cambrian(?)–Ordovician to the Carboniferous. The succession characterizes one of the innermost segments of the wide, north-facing Gondwana platform that once extended from northern South America to the Middle East. Stratigraphic archives of such a proximal segment of the platform remain largely understudied today [Klitzsch, 1994].

Pioneering work occurred at the beginning of the 20th century [Denaeyer, 1924, Fritel, 1924, 1925, with revisions by Lemoigne *et al.*, 1992, Lacroix and Tilho, 1919, Sandford, 1935]. Soon after the Second World War, French geologists of the Borkou-Ennedi-Tibesti mapping project [Bizard *et al.*, 1955, Bonnet *et al.*, 1955, Lessertisseur, 1956, Wacrenier, 1958] recognized that the Palaeozoic of northern Chad is largely comparable to the stratigraphy established by Kilian [1924] for the Central Sahara. The latter comprises fundamentally two geomorphic units: the Cambrian–Ordovician Lower Sandstones resting unconformably on a Pan-African (Proterozoic-earliest Cambrian) basement, and the Silurian–Devonian Upper Sandstones. These units are separated from each other by the intra-Tassilian Trough, which in Algeria and western Libya, includes the lower Silurian graptolitic shales [e.g. the Tanezzuft shales of the Murzuq Basin, Bellini and Massa, 1980]. Publications by German geologists provided details about facies and ichnofacies associations, stratigraphical ages and regional correlation schemes [Klitzsch, 1965, 1968, 1970, Seilacher, 1970, see also Seilacher, 2007]. De Lestang [1965] summarized results acquired by teams of petroleum geologists (SNPA, BRP and PetroPar). This initial set of publications is the basis for subsequent geologi-

cal and hydrogeological synoptic works [Deynoux *et al.*, 1985, MEH (Ministère de l'Élevage et de l'Hydraulique), 2015, Schneider and Wolff, 1992, Trompette, 1983; Figure 2]. Sedimentary logs representative of the Ennedi stratigraphy in Sudan were provided by Klitzsch *et al.* [1993]. On a larger scale, and as the Palaeozoic of the Ennedi Plateau belongs to the Erdi Basin, which is itself the Chadian southern counterpart of the Kufrah Basin [Bellini and Massa, 1980, Bellini *et al.*, 1991, Turner, 1980], comparisons can be made with the adjacent Libyan outcrops and subsurface data, where renewed petroleum exploration has taken place during the last two decades. Le Hérisse *et al.* [2013] and Page *et al.* [2013] described latest Ordovician-earliest Silurian paly-nomorphs and graptolites, respectively. The latter provided some constraints on the age of the fine-grained deposits responsible for the softer “Intra-Tassilian Trough” identified close to the Chad–Libya boundary, which, however, appear older than the classical lower Silurian graptolitic shales. Last, it has been shown that the lower Palaeozoic succession was truncated by a system of late Palaeozoic (early Carboniferous?) ice-stream troughs [Le Heron, 2018, Le Heron *et al.*, 2022].

This paper presents preliminary results of the Lower Sandstones exposed near the town of Fada (Figures 1 and 3), the erosion of which contribute to the magnificent landscapes of the *Ennedi Natural and Cultural Reserve* (<https://www.africanparks.org/the-parks/ennedi>). Relationships with the underlying basement, depositional facies and a refined stratigraphic scheme, as well as the end-Ordovician glacial/post-glacial record will be successively described, completing earlier studies. Finally, the significance of this Ordovician proximal sedimentary record in the overall stratigraphic framework of the Gondwanan cratonic platform is highlighted.

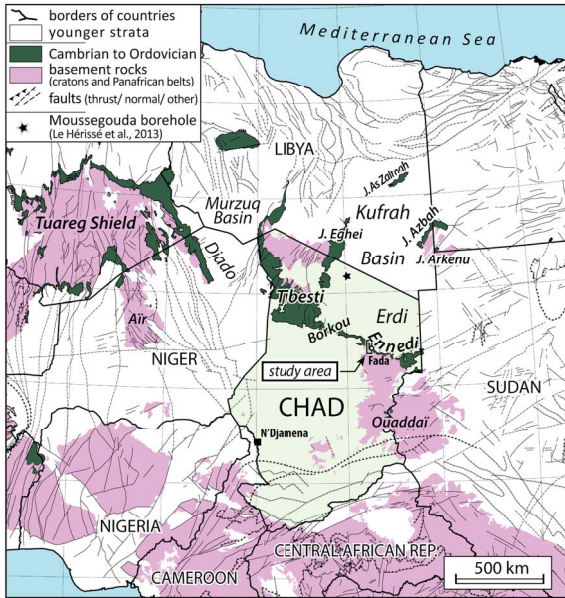


Figure 1. Location map. South of the Kufrah Basin, The Ennedi Plateau (NE Chad) exposes the southern margin of the Erdi Basin [modified from BRGM, 2004, De Lestang, 1965, Le Heron *et al.*, 2015].

2. Lithostratigraphy of the Lower Sandstones and former age attributions

Initially understood as Silurian strata due to the early discovery of *Harlania* ichnofacies [Lacroix and Tilho, 1919; at a time when the Ordovician System was not yet established as such], the Lower Sandstones of the Ennedi were later considered as Ordovician [Wacrenier, 1958]. Afterwards, De Lestang [1965] briefly described the succession distinguished from the base to the top: (i) the formation of the “Basal Sandstone of the Ennedi”, which he assigned to the Cambrian; (ii) a Silurian Wadi Djoua Fm. characterized by ferruginous sandstones; (iii) a fine-grained interval (shales and sandstones, subordinate limestones), referred in northern Chad to as the Bedo Fm., and (iv) the Silurian–Lower Devonian sandstones constituting the Gouring Fm.; the four subdivisions together defining the Borkou Group, the two former corresponding more specifically to the Lower Sandstones (Figure 2). The Lower Sandstones, however, have since been generally assigned to a Cambrian–Ordovician time interval [Trompette, 1983].

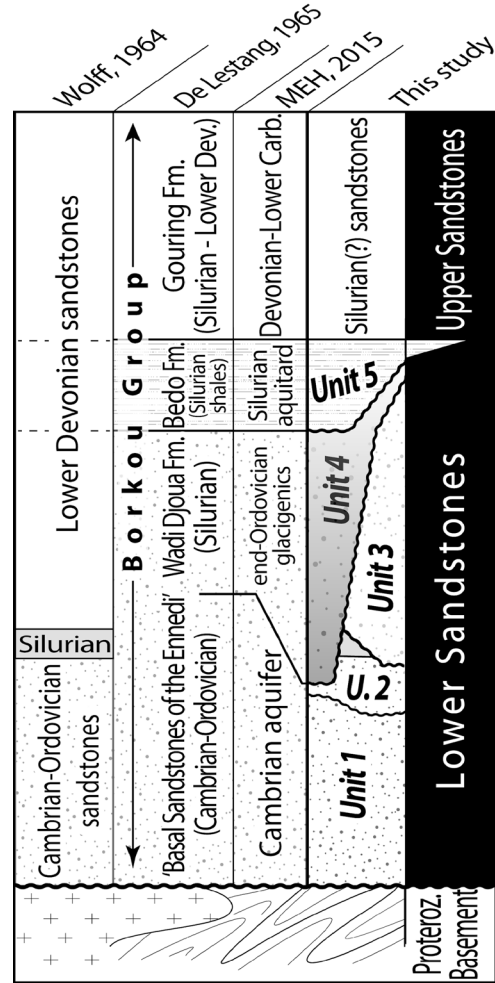


Figure 2. Lithostratigraphic framework and age attributions for the Lower Sandstones of the Ennedi. Four allostratigraphic units (units 1–4) have been identified in the present work. The Bedo Formation (Fm.) (Unit 5), which separates the Lower from the Upper Sandstones, behaves morphologically like the Lower Silurian shaly units across the Sahara, although it may mainly correspond to an uppermost Ordovician interval. Note that another shaly interval positioned at a lower level within the Lower Sandstones (upper part of Unit 2) has been considered Silurian in some previous works.

As noted by Schneider and Wolff [1992], the previous assignment to the Silurian of a shaly horizon responsible for a cuesta mimicking the intra-Tassilian

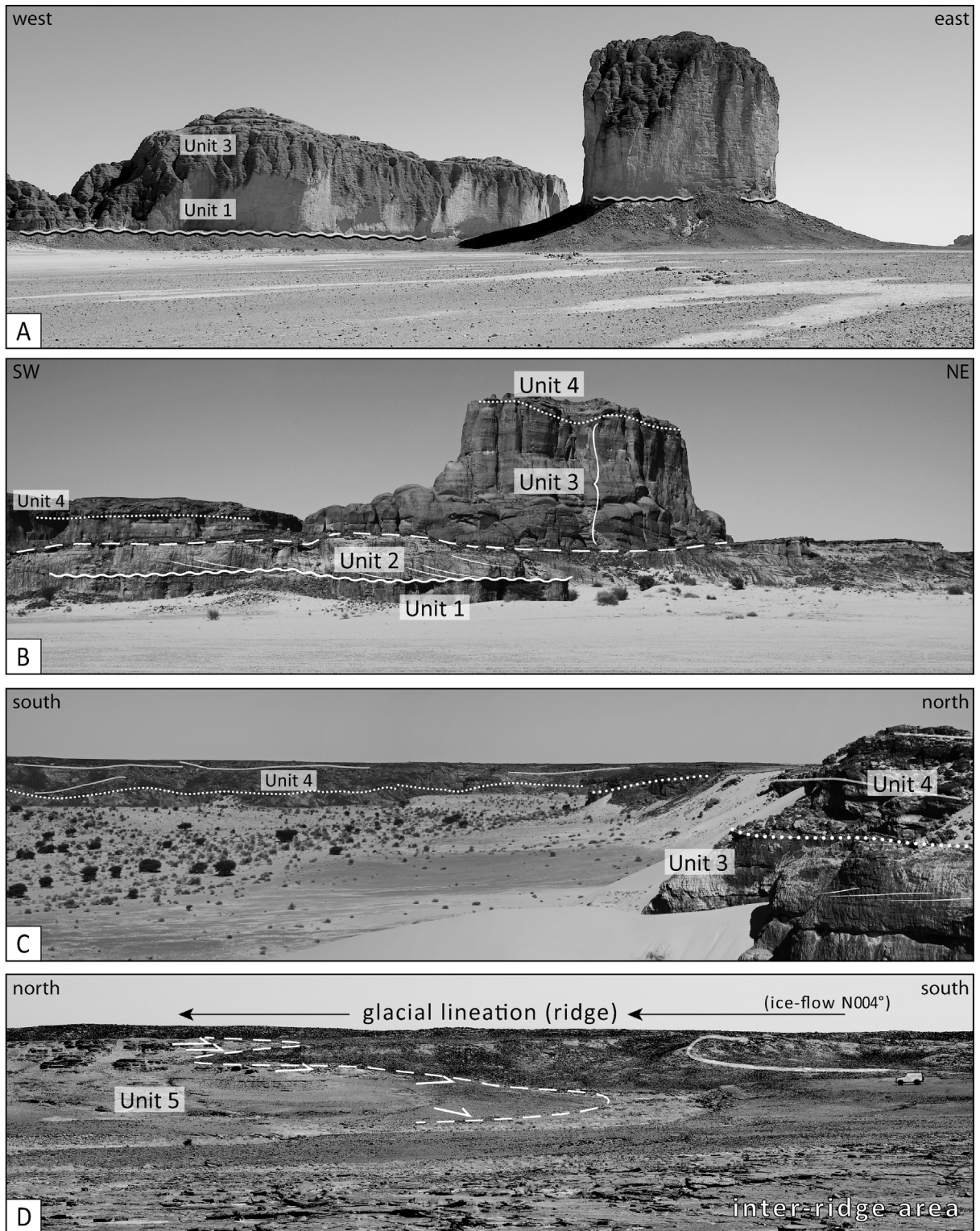


Figure 3. Caption continued on next page.

Figure 3. (cont.) Panoramas showing the characteristics of the Lower Sandstones, from south to north of the study area. (A) The essentially flat lower contact over Proterozoic basement rocks (Deli-Borototou mountain, *c.* 300 m high). In this particular area, Unit 1 and Unit 3 amalgamate, and further amalgamation with the Upper Sandstones is suspected (see Figure 4). (B) Erosional lower bounding contacts of Unit 2 on Unit 1 (wavy line), and of Unit 4 on Unit 3 (dotted line). The fine-grained interval at the top of Unit 2 (not seen) controls the development of a cuesta and mesa. Large-scale cross-strata in Unit 2 are highlighted (Bogaro area; Unit 3 thickness: *c.* 80 m; see also Figure 5C). (C) The regional-scale erosional and undulating contact (dotted line) at the base of the glaciation-related Unit 4 (Wadi Djoua Fm.), which itself includes several, second-order, internal erosional surfaces (see Figure 4). (D) The onlap relationships (white half-arrows) between the fine-grained deposits characterizing Unit 5 (Bedo Fm.) and a late glacial erosion surface, which is partially exhumed and displaying relics of >20 m-high glacial lineations (see also Figure 7D). The latter are locally reworked by iceberg furrows (highlighted cirque behind the 4 × 4; shown with a star in Figure 7D; Wadi Torbo, Bideyat area).

Trough, but positioned well below the Bedo Fm., leads to a pertaining confusion in geological maps where a significant part of the Lower Sandstones is thus wrongly included in the Silurian–Devonian Upper Sandstones [e.g., Wolff, 1964; Figure 2]. The most recent hydrological mapping project [MEH (Ministère de l'Élevage et de l'Hydraulique), 2015] recognizes the true position of the Bedo Fm. in the Ennedi [see also Klitzsch, 1965, De Lestang, 1965] and in addition identifies right below an underlying conglomeratic unit that they attributed to an end-Ordovician “tillite”. They however consider the Lower Sandstones below the latter as a Cambrian rock unit, which suggests preglacial Ordovician strata would be lacking in the Ennedi (Figure 2). However, an essentially Ordovician age will be argued in what follows for the Lower Sandstones of the Ennedi.

3. The Lower Sandstones in the Fada area: overview and basal bounding surface

3.1. Study area and large-scale architecture

The Lower Sandstones have been investigated from the Deli-Borototou mountain (basal contact) to the Wadi Torbo area (upper contact), 40 km to the south and 35 km to the WNW of the town of Fada, respectively (Figure 1 and Table 1). The corresponding submeridian outcrop band dips slightly to the NE and consequently exposes progressively younger strata from south to north. Sandstones emerge from the sand cover in the form of either poorly accessible escarpments, mesa and pinnacles or, alternatively, as extensive horizontal exhumation surfaces (Figure 3).

In the Fada area, the total thickness of the Lower Sandstones is *c.* 200 m. Here we successively describe four allostratigraphic units (Units 1–4) bounded by erosion surfaces from the base to the top of the Lower Sandstones (Figure 4). In addition, an overlying Unit 5 is described, which essentially corresponds to the shaly Bedo Fm. of De Lestang [1965]. Indeed, the Bedo Fm. is still potentially uppermost Ordovician in age, provided the correlation with the fine-grained interval studied by Le Hérissé *et al.* [2013] and Page *et al.* [2013] is valid.

Note that the deep, glaciation-related, erosion surface marking the base of Unit 4 may locally cut down into Unit 2, in which case, Unit 3 is absent. Also, Unit 2 is absent (or at least non-individualized) in places, and in these cases, Unit 3 amalgamates with Unit 1 along a poorly identifiable contact. Unit 4 and Unit 5 are locally very thin or absent, where Silurian sandstones of the Gouring Fm. rest almost directly onto, or locally amalgamate with Unit 3 (Figure 4). This architecture probably explains the various stratigraphic assignments proposed in the literature (Figure 2). Additional work is needed to specify at the regional scale the maximum erosion depths for each of the erosional boundaries.

3.2. The basement-cover unconformity

The Lower Sandstones rest in sharp contact with the underlying Pan-African basement comprising granites, gneiss or quartzite ridges [Saharan metacraton of Liégeois *et al.*, 2013]. The geometry of the contact in the Deli area is essentially horizontal at the km-scale (Figure 3A), but in places displays an undulating surface comprising up to 5 m-deep depressions.

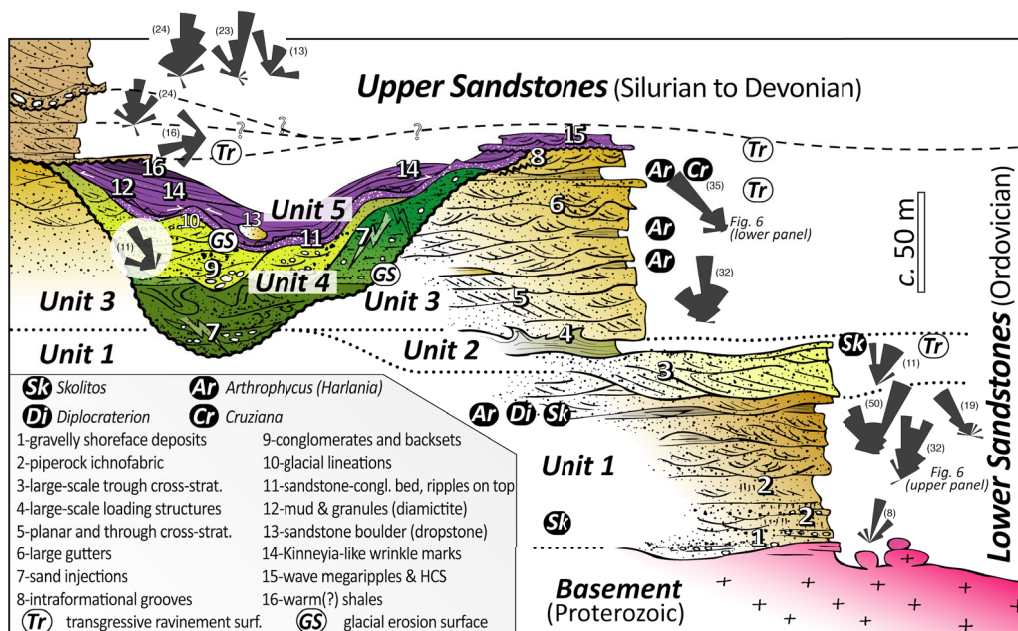


Figure 4. Overview of the Lower Sandstones stratigraphy, which, above the lower unconformity including a lateritic profile, is subdivided in four erosion-based units (see also Figures 2 and 3). Unit 4 includes several erosion-based units and is awaiting a more detailed framework. Note repeated amalgamations where intervening units could not be identified due to non-deposition (interfluve areas) or their subsequent erosion (cf. Figure 3A). Unit 5 (purple, Bedo Fm.) does not belong to the Lower Sandstones, but may be essentially very latest Ordovician in age, with Lower Silurian strata possibly preserved only toward the top of the unit (“warm” shales, Figure 5F). Geographic coordinates in Table 1.

Here, the sandstone cover shows gentle onlapping relationships. Consisting of granites, the uppermost basement rocks show a well-defined tripartite succession, from base to top; massive, weathered granites, outcropping as m-scale balls; deeply weathered granites, corresponding to a soft, up to 5 m-thick, whitish, clayey horizon; a “stratified” unit, 3–8 m in thickness made up of weathered and more or less rubified granites. At the basement-cover contact, a pebbly lag shows dispersed quartzite cobbles up to 15 cm in diameter.

Interpretation. A thick weathering profile had developed before the onset of sand accumulation, and has been largely preserved. Part of the rubefaction is however suspected to correspond to a later diagenetic imprint. Nevertheless, the nature of the pebbles underlining the basement-cover contact—only quartzite, no granite or gneiss—also substantiates a former, intense, weathering phase that left behind only clasts representative of the more resistant

lithologies. Depressions might indicate an intervening erosion phase developing a paleo-drainage network after weathering but prior to deposition of the quartz-rich sediment cover. The weathering profile is reminiscent of the lateritic profile described by Germann *et al.* [1993] in NW Sudan.

4. The Lower Sandstones: preglacial units 1–3

4.1. Unit 1 (lower part of the Lower Sandstones, Figures 3, 5 and 6)

The cliff-forming Unit 1 is nearly 80 m thick. It shows a sandstone-dominated, mainly coarse-grained and cross-stratified succession. At the very base of Unit 1, a specific 3–6 m-thick subunit is individualized. The latter is characterized by alternating gravel conglomerates—which are relatively well sorted (no 1 in Figure 4)—, and coarse-grained sandstones either intensively bioturbated (*Skolithos* piperock

Table 1. Coordinates of observation points where features 1–16 positioned in Figure 4 were best observed. Coordinates of Figure 6 and 5D are also given

Point	Place names		Latitude (N)	Longitude (E)	Related figure
1	Deli	Gravelly-shoreface deposits	16°49.34'	21°27.24'	3A
2	Deli	Piperock ichnofabric	16°49.30'	21°27.27'	3A
3	Bogaro	Large-scale cross-stratification	16°04.88'	21°16.73'	3B, 5B
4	Bogaro	Large-scale loading structure	16°04.94'	21°16.42'	
5	Bogaro	Planar and through cross-stratification	16°05.00'	21°16.43'	5C
6	Hagoto	Large gutters	17°18.85'	21°22.90'	
7	Hagoto	Sand injections	17°16.55'	21°25.18'	
8	Wadi Torbo	Intraformational grooves	17°20.60'	21°19.14'	
9	Fada pass	Conglomerates and backsets	17°12.79'	21°26.06'	7E
10	Wadi Torbo	Glacial lineations	17°19→22'	21°18→19'	3D, 7A-B-D
11	Wadi Torbo	Sandstone-congl. bed, ripples on top	17°20.15'	21°18.78'	
12	Wadi Torbo	Mud and granules (diamictite)	17°18.54'	21°21.06'	
13	Wadi Torbo	Sandstone boulder (dropstone)	17°19.56'	21°19.16'	
14	Wadi Torbo	Kinneyia-like wrinkle marks	17°18.33'	21°21.03'	5E
15	Hagoto	Wave megaripples & HCS	17°15.79'	21°27.08'	
16	Wadi Torbo	Warm(?) shales	17°21.28'	21°19.88'	5F
	Hagoto	<i>Cruziana</i> and <i>Arthropycus</i>	17°18.95'	21°23.06'	5D
	Baki	Large-scale through cross-stratification	~17°02.44'	21°16.58'	6 (upper)
	Hagoto	Large-scale through cross-stratification	~17°19.08'	21°23.69'	6 (lower)

ichnofabric) or made up of 5–20 cm-high, distinctively fining-up, cross-stratified sets showing low-angle laminae. Some of these horizons are radioactive [high U and Th contents; Schneider and Wolff, 1992]. Above, coarse to very coarse sandstones with prevailing cross-stratification are dominant. Active burrowing is still noted, with some recurrences of piperock ichnofabric. *Cruziana* tracks, reported from the base of the succession by Wacrenier [1958] have not been found.

In the middle part of Unit 1, larger-scale and coarser-grained trough cross-strata show less burrowing, with dispersed, rare, or absent(?) *Diplocraterion*. The troughs are consistently oriented toward the NNE (rose diagrams, Figure 4; satellite imagery, Figure 6).

In the upper part of Unit 1 (Figure 4), sandstones show a cyclic pattern corresponding to the superimposition of 3–6 m-thick, erosion-based, fining-up packages. Large (1–2 m high) trough cross-bedded sets of very coarse to granular sandstones, including quartz gravels and showing in place com-

pound cross-stratification and overturned laminae, evolve upwards in progressively thinner and finer cross-bedded coarse-grained sandstones. The latter show frequent to intense, vertical to oblique burrowing. *Skolithos* may form dense networks and *Diplocraterion* is frequent. Cross-stratified sandstones are truncated by a gravel lag or a distinctive, cm- to dm-thick, fining-up granular sandstone bed. Above, but only if preserved beneath the overlying erosion surface marking the base of the next package, is a finer-grained interval. Such intervals, at the origin of discontinuous morphological reentrants that extend at the 100 m scale, correspond either to a single, 5–30 cm-thick horizon of silty very fine-grained sandstone or to a m-thick succession of siltstones and fine- to coarse-grained sandstones (Figure 5A). The latter can be thin-bedded (<5 cm), cross-stratified or entirely rippled, with in places a distinctive lenticular bedding and undulating erosional contacts. Fine-grained, mm-thick laminae may underline some of the oblique laminae and reactivation surfaces are frequently observed.

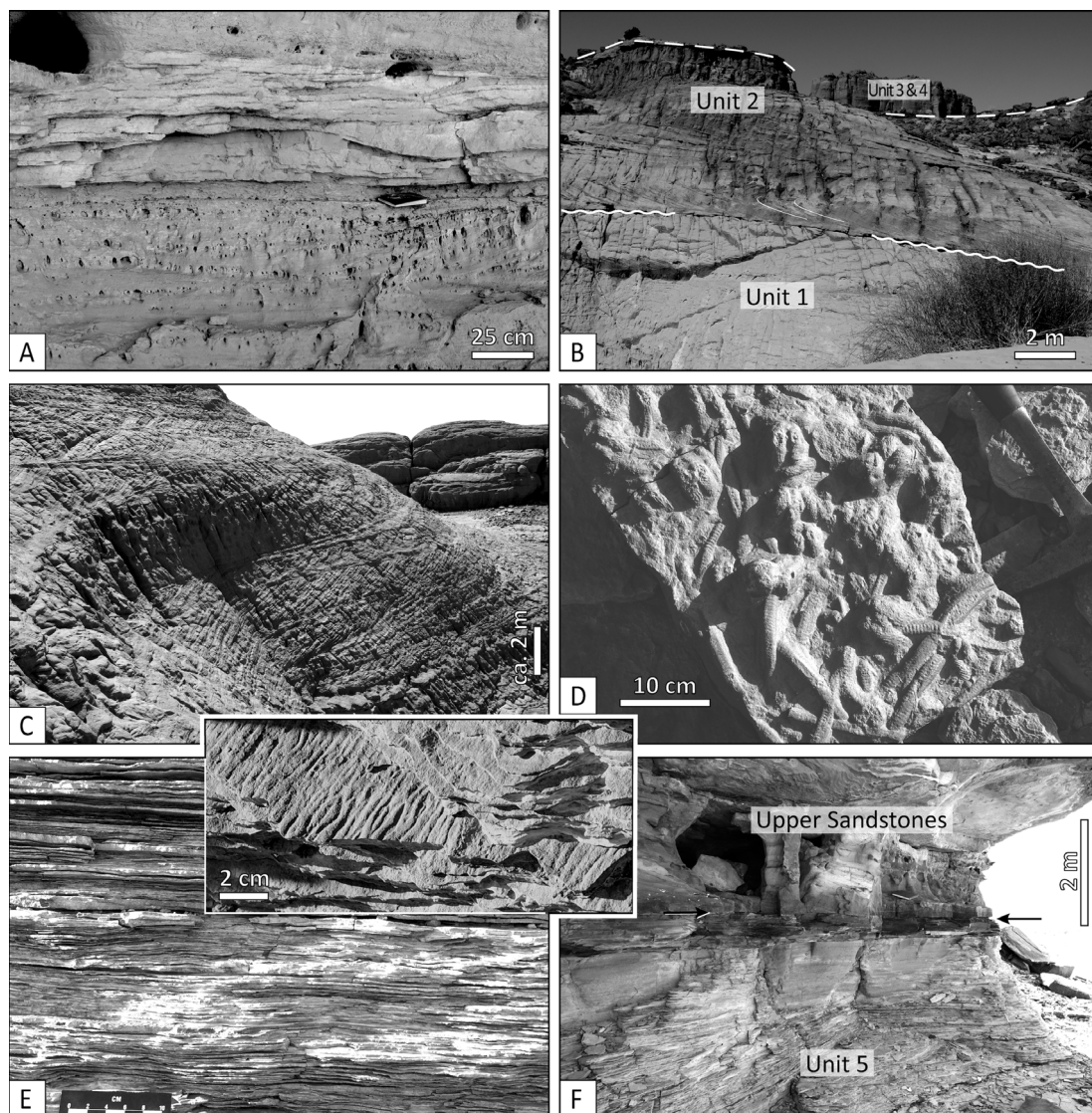


Figure 5. Depositional facies in the Lower Sandstones and Bedo Fm. (A) A tide-influenced interval (undulated bedding, heterolitics, *Arthropycus*) sandwiched between cross-stratified deposits showing a top-truncated coarse-grained, bioturbated sandstone (lower part) and an overlying erosion-based very coarse-grained, non-bioturbated sandstone (upper part of Unit 1). (B) Detailed view of Figure 3B, shows the erosion-based Unit 2, large-scale and rhythmic cross-stratification (including overturned laminae; Figure 4/3) and the truncation by a transgressive ravinement surface in its upper part (dashed line). (C) Mixed planar and trough cross-stratification (non-bioturbated) in the lower part of Unit 3 (Figure 4/5). (D) Associated *Arthropycus* and *Cruziana* (mainly *Rusophycus* forms), characterizing a singular fine-grained interbed in the upper part of Unit 3. (E) Rhythmically laminated fine-grained sandstones in the lower part of Unit 5 (Figure 4/14). Layer tops typically show *Kinneyia*-like wrinkle marks (onset). (F) The topmost part of Unit 5 made up of shales, lacking sandstone horizons [warm? shales of Meinhold *et al.*, 2016; Figure 4/16], and truncated by a distinctive coarse-grained sandstone bed (black arrow). Seen above is the base of Upper Sandstones (Silurian to Devonian); geographic coordinates in Table 1.

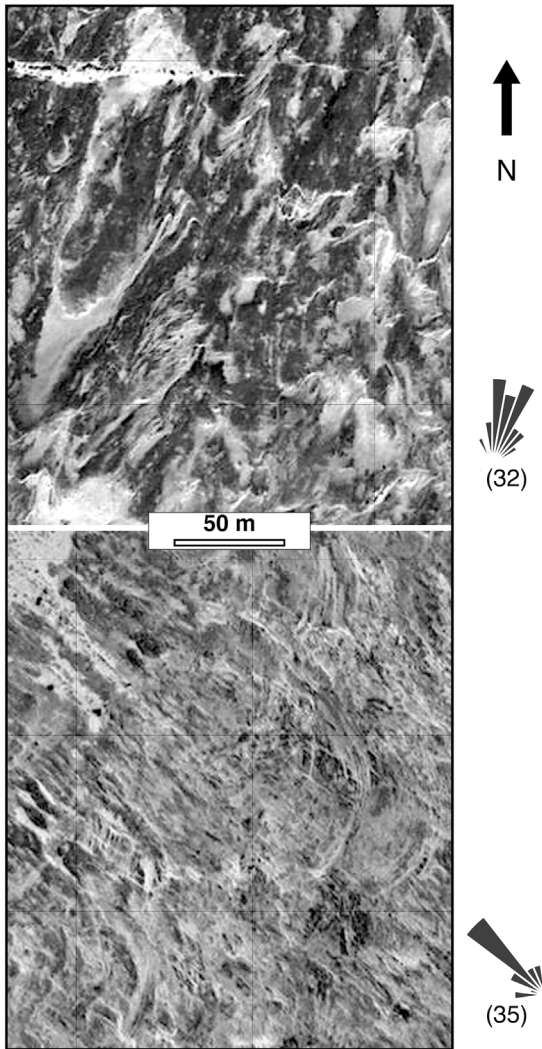


Figure 6. Satellite images illustrating two types of large-scale trough cross-stratification in the Unit 1 (upper panel) and Unit 3 (lower panel) of the Lower Sandstones. Rose diagrams of cross-stratal dips were made from areas near (<3 km) the views (stratigraphic locations in Figure 4; geographic coordinates in Table 1).

Such intervals show *Harlania* (= *Arthropycus*) everywhere, a favorable heterolithic sedimentation has developed.

Interpretation. Unit 1 is dominated by coastal deposits, at the transition from fluvial to nearshore environments. Basal conglomerates and related sandstones are more specifically understood as gravelly

shoreface deposits [Dashtgard *et al.*, 2009, Hart and Plint, 1995] in an overall transgressive setting; the latter being also indicated by active burrowing and the concentration of radioactive minerals [Dabard *et al.*, 2015].

In the middle part of Unit 1, the fluvial signature prevails but then decreases again upwards in parallel with occurrences of thin but apparent shallow-marine incursions showing a tidal signature (gravel lag indicating ravinement processes, drapes of fine-grained sediments, lenticular bedding, intense bioturbation). Rippled sandstones in these finer-grained intervals are tentatively interpreted as estuarine mouth-bar deposits.

From base to top of Unit 1, a continuously aggrading framework is suggested, nevertheless recording a long-term transgressive/regressive/transgressive evolution. Being described as erosion-based fining upward packages, high-frequency “transgressive” cycles are suggested as building blocks of the upper part of Unit 1. However, the cyclic pattern might alternatively represent the superimposition of “highstand” regressive (large cross-bedded sets) then transgressive (bioturbated cross-bedded sandstones and overlying tide-influenced intervals) deposits. In the latter case, the erosion surface at the base of any individual package would correspond to channel bases rather than to a regional-wide subaerial erosion surface. This may explain the lack of indication for weathering in between individual cycles. An overall marginal-marine depositional environment, comprising fluvial, bay-head delta and estuarine deposits, is inferred on the basis of both primary sedimentary structures and associated bioturbation [Buatois and Mángano, 2011]. Whether the high-frequency alternation of fluvial-dominated and nearshore deposits represent autogenic or allocyclic processes—e.g. controlled by eustacy or a varying rate of sediment supply—remains to be elucidated [Zecchin *et al.*, 2007].

4.2. Unit 2 (middle part of the Lower Sandstones, Figures 3B and 5B)

Unit 2 is a relatively thin, bipartite succession resting onto an erosion surface truncating Unit 1. The main part of Unit 2 is a <15-m thick, fining and thinning-upward sequence of cross-bedded sandstones. At the base, sandstones are organized in

large-scale trough cross-stratified sets, up to 4 m in thickness (Figure 4/3). A rhythmic pattern can be identified when <5 cm-thick granular laminae repeatedly alternate with up to 40 cm-thick medium- to coarse-grained sandstones including dispersed gravels. Upward, cross-bedded sandstones typically show thinner, m-thick sets. Bioturbation has not been observed. Cross-bedded sandstones are sharply truncated by an essentially flat erosion surface underlined by a gravel lag including some quartz pebbles at the base of a 10–20 cm thick distinctive graded sandstone bed, including poorly defined burrows (*Skolithos*?). The uppermost part of Unit 2 consists of 1 m-thick, highly ferruginized siltstone, from which primary structures have been erased. Siltstones are deformed by decameter-scale load casts affecting the immediately overlying sandstones of Unit 3.

Interpretation. A fluvial origin could be proposed from the lower cross-bedded sandstones, even if the rhythmic organization might suggest a tidal cyclicity; the advance of coarse-grained, bar-top megaripple trains would, however, also produce such a rhythmic pattern without any tidal signal. The truncation and the related coarse-grained bed is interpreted as a marine ravinement surface overlain by a horizon of “basal transgressive sand”. The overlying siltstones may reflect either offshore deposition or fine-grained sedimentation in the middle, low-energy segment of an estuarine system. However, diagenetic ferruginization precludes a more advanced interpretation.

Depositional environments of Unit 2 do not actually depart from the ones identified in the upper part of Unit 1 (and overlying Unit 3). A deeper basal erosion surface, larger cross-bedded sets and a (residual) 1 m-thick siltstone interval each appear as exaggerated versions of all the features characterizing cyclical packages described in the upper part of Unit 1. A larger-scale pulsation (higher amplitude? higher rates of variation?) of the factors controlling coastal sedimentary dynamics in the Ennedi is thus inferred at this level.

4.3. Unit 3 (upper part of the Lower Sandstones, Figures 3B, 5 and 6)

Like Unit 1, Unit 3 is a cliff-forming, up to 100 m-thick succession of essentially cross-bedded sandstones. Where Unit 2 ends with siltstones, the basal

erosional contact of Unit 3 is underlined by a distinctive blue-grey, coarse-grained gravelly sandstone showing large-scale load structures (Figure 4/4). Unit 3 is characterized by coarse- to very coarse-grained cross-stratified sandstone beds with dispersed gravels. Mixed large trough and planar cross-stratification are identified (Figure 5C; Figure 4/5). Neither bioturbation nor fine-grained intervals have been observed in the lower 30 m of Unit 3 but fine-grained intervals including *Arthropycus* reappear upwards (only observed in scree deposits).

The upper part of Unit 3 is only exposed in a few preserved interfluvial areas beneath the deep erosion surface characterizing the base of the overlying, erosion-based Unit 4 (Figure 4). Medium- to coarse-grained cross-strata are observed, corresponding to 10 m wide troughs; the larger being distinguishable from satellite images and evidencing a NW-oriented dispersal pattern (Figure 6). As in Unit 1, cross-bedded sets are organized in fining- and thinning-up packages. In contrast of the ones of Unit 1, some of the intervening bounding erosion surfaces are deeply (>2 m) incised with pebbly lags and steep margins suggesting large-scale gutters. In addition, cross-bedded sandstones show less bioturbation than those of Unit 1. A dm-thick, distinctively well-sorted, coarse-grained sandstone bed underlines the boundary in places between cross-bedded sandstones and the overlying fine-grained interval. In the latter, bioturbation includes small Cruzianids, generally *Rusophycus* forms, rarer *Cruziana* forms, which are associated, or not, with *Arthropycus* showing well-defined annulations (Figure 5D).

Interpretation. A large-scale transgressive trend is recorded from the base to the top of Unit 3, as evidenced by the greater development of fluvial deposits in the lower part of the unit (including, frequent planar cross-stratification), the overall fining-up trend and ichnofacies changes (no bioturbation, then *Arthropycus*, then *Arthropycus* and *Cruziana*). Though it remains to be confirmed, a lesser development of the *Skolithos* ichnofacies seems to be noted in the whole of Unit 3 when compared to Unit 1. Well-sorted sandstone beds separating fluvial sandstones from marine, bioturbated intervals are interpreted as the signature of marine ravinement processes, possibly by tidal currents.

5. The record of the end-Ordovician glaciation: units 4–5

5.1. Unit 4 (*uppermost part of the Lower Sandstones, Wadi Djoua Fm., Figures 3D and 7*)

Unit 4 constitutes the infill of several tens of meters deep incisions truncating Unit 3, which are themselves truncated by the erosion surface marking the base of Unit 5 (Figure 4). Unit 4 consists of a variety of massively bedded to cross-stratified sandstones, with various abundance of dispersed pebbles. Sandstones locally grade into conglomerates with dispersed quartzite cobbles up to 20 cm in diameter. Unit 4 is also characterized by discontinuous chaotic outcrops that include cuesta-like slopes at the km-scale, large-scale (>100 m) sigmoidal and channelized sandstones bodies, sand injections in the form of 5–15 cm-thick ferruginous dikes or sills (Figure 4/7), and three-dimensional folding structures affecting 1–3 m thick sandstone horizons.

No paleovalley network has been properly identified so far, yet WNW-EWE-oriented depressions have been inferred, in agreement with paleocurrent trends (Figure 4). In “interfluves” area, Unit 4 is essentially missing and a thin or absent Unit 5 rests directly on Unit 3. Unit 4 is equivalent to the Wadi Djoua Fm. of De Lestang [1965] (Figure 2). However, the author reported dolerite intrusions in this formation, which we did not observe.

A detailed facies succession in Unit 4 remains to be established and only preliminary observations are provided below. The basal bounding surface and immediately overlying sandstones were generally affected by heavy ferruginization, which hinders the identification of probable onlaps relationships. Conglomeratic facies associations appear preferentially in the lower part of the unit, while the initial deposits that directly seal the basal erosion surface of Unit 4 are, in places, made up of relatively well-sorted coarse-grained sandstones. Cut-and-fill structures 0.5–5 m-deep and 5–50 m in width show alternation of well- to faintly-laminated coarse-grained sandstones and clast- or matrix-supported pebbly conglomerates with dispersed cobbles. Undulating laminae, backsets and pebbly lags are also noted (Figure 7E). Clasts are subangular to subrounded and clast lithologies are both extrabasinal (vein quartz, fine-grained, granular or gravelly quartzites, rare cherts) and intrabasinal (sandstones and siltstones).

The latter may constitute intraformational conglomerates, with siltstones intraclasts up to 60 cm-long and a poor representation of extra-basinal clasts. Such conglomeratic deposits progressively grade upward in coarse- to medium-grained, trough cross-stratified sandstones. Together they constitute tens of m-thick, fining- and thinning-upward sandstone sequences. Such successions include in places m-thick intervals of thinly bedded, fine- to medium-grained sandstones with rippled tops and small-scale load casts, and are locally truncated by sandstone beds characterized by abundant rip-up clasts at their top. Though the former have a limited extent (10s to 100s of meters), the latter constitute extensive, horizontally stratified marker beds at the km- to 10 km-scale, which sharply truncate the underlying chaotic and locally deformed sandstones forming the main infill of the incisions.

Interpretation. The deep basal erosion surface, the overall chaotic organization, the conglomerates with numerous extra-basinal clasts, and the stratigraphic position immediately beneath a well-identified glacial surface (see Unit 5) permit Unit 4 to be interpreted as a glaciation-related depositional unit. Similar facies associations including subglacial sedimentary injections and large-scale channels have been documented in the Felar-Felar Fm. of the Djado Basin [north Niger; Denis *et al.*, 2010] and in the Mamuniyat Fm. of the Kufrah Basin [southeastern Libya; Le Heron *et al.*, 2010, 2015]. Large intrabasinal clasts and bedding/lamination patterns that indicate supercritical flow processes noticeably depart from the other fluvial deposits described in units 1–3 and also support the glacial interpretation [Girard *et al.*, 2015, Lang *et al.*, 2021]. No striated pebbles have been, however, identified so far. This is usually the case in bedload dominated deposits, where tractive processes rapidly erase potential glacial striae. Cryptic intraformational glacial erosion surfaces might be suspected owing to the occurrence of large-scale folds and massive sandstones, which could represent deformation tills especially where they are intruded by networks of sandstone dikes [Ravier and Buoncristiani, 2018]. Channelized and sigmoidal sandstone bodies at the 100 m scale in one hand, and intervening finer-grained intervals including flaser bedding and load casts are reminiscent of end-Ordovician meander belts [Rubino *et al.*,

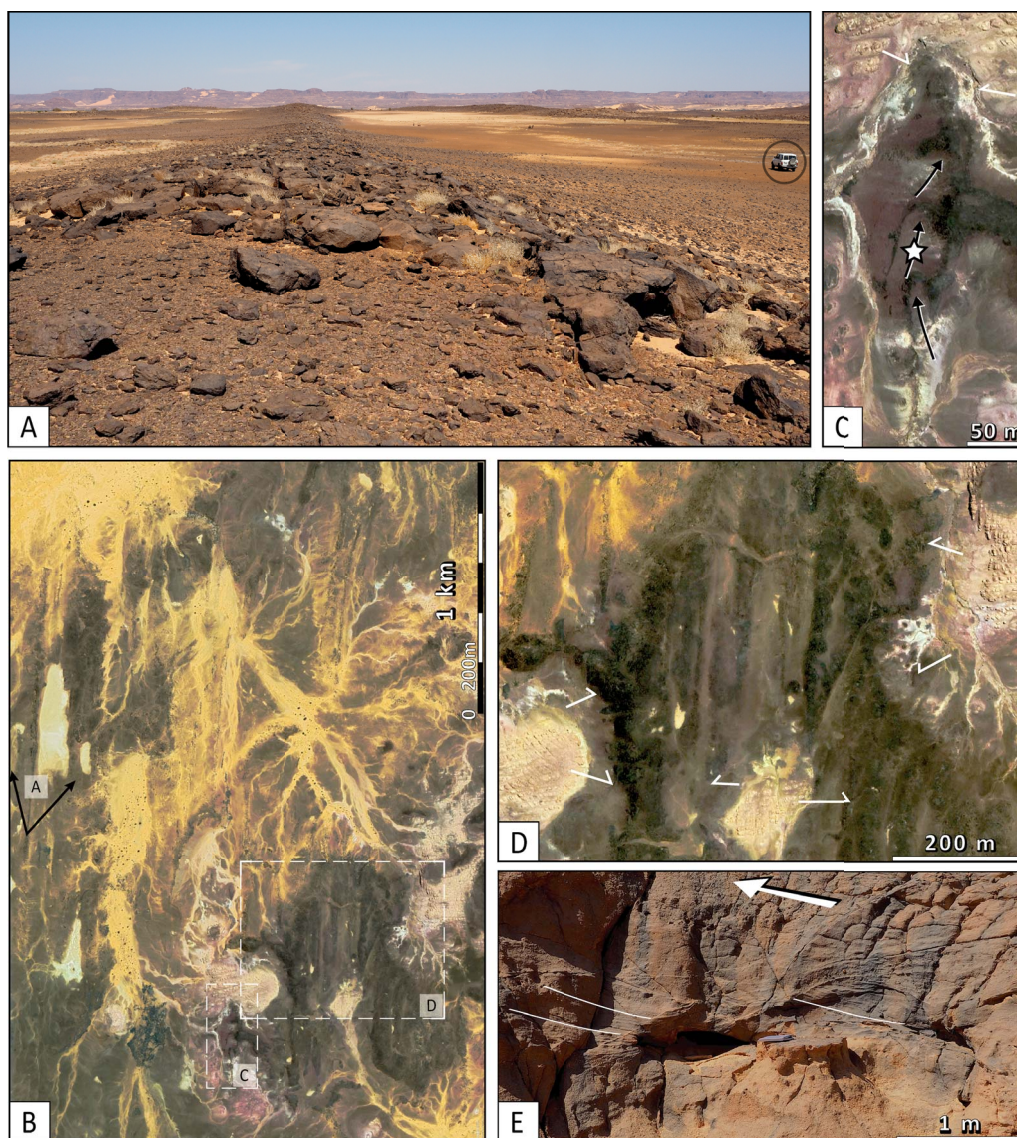


Figure 7. The end-Ordovician glacial record in the Ennedi. (A) A streamlined glacial lineation of the corrugation field at the lower part of Unit 5 (location in Figure 7C; 4×4 for scale). (B) Satellite view (Google Earth image) of the corrugation field of the Wadi Torbo area, showing a north-oriented ($N004^\circ$) ridge-and-trough topography and superimposed second-order ridges (image center at $17.33^\circ N$, $21.33^\circ E$). Close-up views of (C) superimposed furrows interpreted as iceberg ploughmarks reworking a glacial lineation (star: see Figure 3D) and (D) second-order ridges onto a larger-scale ridge. White half-arrows highlight the onlaps of Unit 5 deposits (light colored residual patches) onto the sandstones (dark) corresponding to glacial lineations of higher elevation (Figure 4/13). (E) Pebbly sandstones of Unit 4 recording prevailing upper-flow depositional processes (Figure 4/9; faint and/or undulated laminations, backsets, intrabasinal cobbles). Highlighted backsets (white), flow almost from right to left (black arrow).

2003]. Fluvio-glacial dynamics in an ice-marginal context is therefore predominant in Unit 4, though intervening subglacial deformation events are also suspected. Sandstone marker beds are tentatively interpreted as subordinate transgressive horizons, a better identification of which should help to establish a more detailed stratigraphic succession in Unit 4.

5.2. Unit 5 (*Bedo Formation, Figures 5 and 7*)

Unit 5 was observed close to Wadi Torbo, where it is discontinuously but frequently exposed over a c. 100 km² area. Elsewhere, Unit 5 might not be individualized (for instance, south of Fada; Figure 3A), in which case the identification of the contact between the amalgamated Lower and Upper Sandstones is challenging (Figure 4). Unit 5 is mainly comprised of a fine-grained succession including an intervening sandstone interval. This succession onlaps at the outcrop scale a corrugated surface shaping the ferruginized top of an underlying sandstone (white arrows in Figure 7D).

The corrugated surface is shaped in parallel, north-south ($N 04 \pm 1^\circ$) oriented ridges and troughs (Figure 7B). Wavelengths and amplitudes are in the 250–500 m and 5–30 m ranges, respectively. Some of the corrugations can be followed individually on distances >3 km (Figure 7A). Rectilinear, smaller-scale (<10 m wide), second-order and parallel sandstone ridges show higher elongation ratios (Figure 7D). Beneath the corrugated surface, the fine-grained succession is almost everywhere underlined by a continuous, <1 m-thick, sandstone/conglomerate bed, which therefore appears to be also conformable to an underlying, undulating erosion surface (Figure 4/11). Small-scale extensional step fractures are ubiquitous in the coarse-grained bed. Only one isolated (displaced) stone bearing striae was discovered. In a single outcrop, N–S-oriented poorly defined intraformational grooves have been observed in the underlying Unit 3, one meter below the Unit 5/Unit 3 contact (Figure 4/8). The sandstone/conglomerate bed locally shows a pebbly lag at its upper surface and is systematically capped by trains of current ripples atop a cm-thick veneer. The lag and the ripple trains mold m- to several m-scale gullies or depressions that add local complexity to the otherwise relatively smooth corrugated surface. Satellite images reveal

that the corrugation field is disrupted by 25–60 m wide furrows with lateral and frontal berms. Their orientation ranges from N350 to N30, the frontal berm being positioned to the north (Figure 7C).

The fine-grained succession above the corrugation field starts with a thin (5–20 cm) muddy sandstone or siltstone horizon including in places isolated quartz granules. Above, and constituting the main part of Unit 5, are laminated, whitish, clayed siltstones including various proportions of fine-grained, thin-bedded sandstone (mm-scale “laminae” to cm-scale beds). Rhythmical patterns are frequently observed at the 10–20 cm scale. A sandstone boulder has been found isolated in the lower part of Unit 5 (Figure 4/13). A more indurated horizon is in addition usually observed within the soft, fine-grained succession. Its nature and position are variable from place to place, yet usually observed in the lower part of Unit 5 (Figure 4/14). Additional work is needed to confirm it represents a correlatable horizon. In troughs, this horizon may not fundamentally differ from the laminated siltstones (less clayed?) or corresponds to a few m-thick, heretolithic interval characterized by a flaser to lenticular bedding at the mm- to cm-scale. In this case, Kinneyia-like wrinkle marks repeatedly characterize bed tops, showing a consistent N20–N200 orientation (Figure 5E, Figure 4/14). Laterally, where Unit 5 thins out over “paleohighs” (the highest ridges of the corrugation field, i.e., less eroded interfluvial areas), a single coarse-grained sandstone bed, 20–40 cm-thick, relatively but distinctively well-sorted, shows cross-stratification and, in some places, megaripples with preserved primary relief (Figure 4/15). An intraformational conglomerate of rip-up siltstone clasts occasionally underlines the lower bed contact. Above, dm-thick, fine- to medium-grained sandstone beds show in places hummocky cross-stratification (HCS) laminations with parting lineations and/or *Arthropycus*.

The uppermost 2 m of Unit 5 consists of shales (Figures 5F and 4/16). This is the finest-grained interval of Unit 5, enhancing an overall fining-upward pattern. Unit 5 is sharply truncated by a dm-thick gravelly sandstone bed underlining the trough cross-bedded Silurian–Devonian Upper Sandstones. The latter consist of medium- to coarse-grained sandstones including recurrences of siltstone horizons and bioturbations, mainly tracks along horizontal

truncation surfaces. Palaeocurrent trends noticeably deviate to the west or SW (Figure 4). Upper Sandstones grade upwards into coarse- to very coarse-grained deposits, with resumption of palaeocurrent trends toward the north.

Note that due to onlaps onto the corrugation field (Figures 4 and 7), Unit 5 is most often incomplete; Unit 5 being reduced to a few m-thick interval above some of the highest ridges. The lowermost Upper Sandstones have a wider distribution but might also be absent (or undifferentiated?). It has not been established whether their absence is related to non-deposition or to truncation beneath one of the erosion surfaces locally underlined by intrabasinal conglomerates within the lower part of the Upper Sandstones (question marks in Figure 4).

Interpretation. The corrugation field is interpreted as an exhumed a glacial surface. The ridge-and-trough organization and superimposed second-order ridges correspond to streamlined glacial lineations developed over a soft sediment substrate [Deschamps *et al.*, 2013, Le Heron *et al.*, 2022, Moreau *et al.*, 2005]. In this context, the sandstone/conglomerate bed draping the undulating basal truncation likely corresponds to a subglacial fluidized horizon [Denis *et al.*, 2010]. In association with late step fractures and underlying intraformational shear structures such as grooves, this bed highlights a soft-sediment subglacial shear zones [Denis *et al.*, 2010, Girard *et al.*, 2015, Le Heron *et al.*, 2005]. In map view, the Chadian corrugation field resembles assemblages of Quaternary megascale glacial lineations (MSGL), but we note however that: (i) the width of the highly-elongated second-order ridges is noticeably smaller than that of typical MSGL; and (ii) conversely, the height of the first-order corrugations, in excess of 20 m, is largely greater than that of typical MSGL [Spagnolo *et al.*, 2014]. This suggests that the corrugations fit a specific category of glacial lineations, characterized by vertical amplitude greater than 20 m, which can represent forms that are transitional with elongated drumlins [Spagnolo *et al.*, 2014]. It is in addition noted that the core of the corrugations is made up of non- to slightly deformed Unit 3 sandstones. Glacial lineations observed in Wadi Torbo are therefore interpreted as having been formed according to a drumlinization model accounting for the erodent layer hypothesis

[Eyles *et al.*, 2016]. Initial drumlins essentially shaped by carving in pre-existing bed material were progressively reworked into ridges under an accelerating ice stream. Delicate features preserved at the former ice-sediment contact (striae, grooves...) were not preserved due to active and generalized post-glacial winnowing evidenced by ripple trains (tidal pumping close to a retreating grounding line?). Only the largest structures were preserved, such as gullies (formed by gravity-driven instabilities?), furrows and associated berms (ploughmarks of northward-drifting icebergs) or depressions (marks of keels of stranded icebergs?). A potential analogue of this palaeo-landsystem is the Barents Shelf [Newton and Huuse, 2017], considering the intracontinental setting, moderate water depths, MSGL size and cross-cutting relationship with iceberg ploughmarks.

Granules dispersed in siltstones and muddy sandstones draping the glacial lineations most likely record ice-rafting processes. Interestingly, no outwash facies have been identified. Overlying laminated siltstones, lacking ice-rafted clasts—with, however, the notable exception of the sandstone boulder—, lacking storm-generated facies and bioturbation, but showing rhythmical stratification pattern are tentatively interpreted as the settling of hypopycnal meltwater plumes beneath a perennial sea-ice cover [Rüther *et al.*, 2012]. In this case, and accounting for high ice-marginal accumulation rates, the time span represented by the fine-grained Unit 5 might be relatively short [Lucchi *et al.*, 2015]. The sea-ice cover hypothesis is favored relative to an ice-shelf model because of the shallow water depths, the lack of ploughmarks from large tabular icebergs, and the lack of facies marking the retreating calving-line [Smith *et al.*, 2019]. The more conservative interpretation that would simply consider a restricted-marine depositional environment for the Bedo Fm. cannot be ruled out.

In this context, the significance of the intervening coarser-grained interval observed in the lower part of the Bedo Fm. remains unclear: the break-up of the ice cover and restoration of storm processes; temporary re-advance of a grounding line; or transgressive deposits involving a healing phase [Proust *et al.*, 2018]? Over the adjacent highs, HCS beds and megaripples, which are the signature of combined flow deposits in fine- to medium-grained, and coarse-grained sands, respectively [Leckie, 1988], may be the proximal

counterpart of such a healing phase. In any case, wave processes suggest that the Ennedi was affected at that time by relatively open-marine conditions.

The presence of Kinneyia-like structures, most likely linked to microbial mats [Aubineau *et al.*, 2018] suggests a tidal influence with cyclical, microbially-mediated, hydrodynamic instabilities [Thomas *et al.*, 2013]. Crest orientation of the wrinkle marks conforms to submeridian depositional slopes, i.e., dipping almost parallel to the axis of former glacially cut overdeepenings. The development of microbial mats is, however, inconclusive regarding the presence, or the termination, of a sea-ice cover. The overall Unit 5 might be the signature of an ice-front retreat during a glacio-isostatically induced forced regression rather than that of a postglacial transgression [Dietrich *et al.*, 2019]. Only the 2 m-thick shale horizon at the very top of the Bedo Fm. suggests an ultimate, post-glacial sea-level rise.

The truncation of the Bedo Fm. by cross-bedded sandstones showing paleocurrents departing from regional trend (Figure 4) is interpreted as an extensive marine ravinement surface overlain by tidal sandstones. An unconformity and a major hiatus are thus inferred at the top of the late- to postglacial shale-dominated sedimentation. Similar relationships are also been identified in the analogous succession of the southeastern Kufrah Basin [Turner, 1991], which interestingly differs from more continuous regressive trends interpreted from more basinal areas [Lüning *et al.*, 2010, Gindre *et al.*, 2012].

6. Discussion: Significance of the Ennedi record

6.1. The preglacial succession

The study area is located at *c.* 600 km to the SSE and SSW from the closest outcrop belts fringing the Kufrah Basin (Jebel Eghei and Arkenu area, respectively), and at *c.* 900 km to the ESE from the Djado (Niger). Considering NNE- to NW-ward oriented paleocurrents, the study area is representative of one of the innermost segments of the Gondwana shelf (Figures 1 and 4). In the absence of biostratigraphic dating, only attempts at regional correlation can be proposed for the Lower Sandstones of the Ennedi. The latter are comparable in terms of depositional environments to the Lower Sandstones around the

southeastern Tuareg Shield in northern Niger [also known as the Timesgar Sandstone; Joulia, 1959] and SE Algeria [sometimes referred to as the Tin Taharadjeli Member; Beuf *et al.*, 1971]. The stratigraphic assignment of these strata has long been controversial. While a Cambrian to lowermost Ordovician age is often proposed [e.g., Eschard *et al.*, 2010, Greigert and Pougnet, 1967, Perron *et al.*, 2018], the same succession is regarded in the Tassili area as Ordovician by Legrand [1974, 2002], Riché and de Charpal [1998] and Fabre and Kazi-Tani [2005]. The Lower Sandstones of the Ennedi are also similar to successions exposed in the Kufra Basin (fluvial sandstones with overturned cross-stratification, tidal bars, superimposed m-scale cycles, among others; see detailed description by Turner [1980]). Originally attributed to a Cambrian Hasawnah Fm. [Bellini and Massa, 1980], the upper portion of the Lower Sandstones of the Kufrah Basin is today mostly correlated with the shallow-marine, (Lower to?) Middle Ordovician Hawaz Fm. of the Murzuq Basin [Anfray and Rubino, 2003, Gil-Ortiz *et al.*, 2022]. This revision results from the generalized observation in SE Libya of bioturbated sandstones including *Skolithos*, *Arthropycus* and *Cruziana* [Ghienne *et al.*, 2013, Le Heron and Howard, 2012, Lüning *et al.*, 1999, Seilacher *et al.*, 2002]. Of particular interest is a change in sedimentation identified at *c.* 200 m beneath the glacially-related strata in the Jebel As Zalmah area, involving a basal conglomerate, a significant transgressive episode and an overall change from prevailing trough cross-strata below, to planar cross-strata above [Le Heron and Howard, 2012]; all of these features being reminiscent of Unit 2 and its relationships with Unit 1 and Unit 3 of the Ennedi (Figure 4). Unit 2 might also correlate with the middle part of the Lower Sandstones of the Jebel Arkenu [Turner, 1980] and the lower Araske Fm. of the Tibesti [De Lestang, 1965]. It is worth noticing that in the Ennedi, the Lower Sandstones include bioturbated coastal deposits since their very base, lacking the thick fluvial deposits characterizing the lower/lowermost part of the Cambrian–Ordovician succession in the more distal Murzuq Basin [Elicki and Altumi, 2019, Ghienne *et al.*, 2013].

As a working hypothesis, and considering the occurrence of *Cruziana petraea* in the upper part of the Lower Sandstones of the Kufrah Basin [Lüning *et al.*, 1999], an ichnospecies generally characterizing

Middle to Upper Ordovician deposits reported from Chad [Seilacher, 2007] as well, we suggest the following attributions for the pre-glacial Lower Sandstones units on the basis of the recognition of two large transgressive cycles, which we acknowledge as two large Ordovician retrogradations over the platform.

- Unit 1: mostly Lower Ordovician sandstones, which should correspond to the proximal counterpart of sandstones recurrently showing *Cruziana* tracks of the Rugosa group more to the west and north. The lowermost shallow-marine segment might be the only representative, if any, of an (upper?) Cambrian sedimentation, though an earliest Ordovician age is favored.
- Unit 3: upper Lower to Middle Ordovician sandstones, which we correlate with the Floian to Darriwilian retrogradation recorded to the west (In Tahouite Fm. of Algeria; Hawaz Fm. of Libya). The overall “transgressive” trend is highlighted by conditions favorable to a *Cruziana* ichnofacies development at the top of Unit 3.

An alternative correlation might have considered that the two transgressive wedges relate to Early-Middle (Unit 1) and Late (Unit 3) Ordovician cycles; the latter being well known for a major retrogradation over north and west Africa [Gatinskiy *et al.*, 1966, Kichou-Braïk *et al.*, 2006] and the Middle East [MFS O40 of Sharland *et al.*, 2001]. Upper Ordovician strata are, however, most often truncated by the end-Ordovician glacial erosion surfaces, except in the more distal segments of the platform [Ghienne *et al.*, 2010] or in interfluvial areas beneath the end-Ordovician glacial erosion surfaces [Bataller *et al.*, 2019, Gil-Ortiz *et al.*, 2022]. Though these alternative attributions cannot be excluded, the change in paleocurrent trends noted from Unit 1 to Unit 3 (Figure 4) is reminiscent of the change in drainage patterns that occurred in an Early to Middle Ordovician time interval [Beuf *et al.*, 1971], favoring the Early to Middle Ordovician age attribution considering the Lower Sandstones of the Ennedi.

The status of the intervening Unit 2 is puzzling. It might be understood as an ultimate short-term “parasequence” concluding the long-term transgressive stacking pattern initiated with Unit 1 and trun-

cated by Unit 3, which represents an overlying transgressive wedge similar to Unit 1. This interpretation would suggest the Unit 3/Unit 2 contact corresponds to a relatively long stratigraphic hiatus, which should correspond to the development of highstand to lowstand conditions in more distal areas. It however does not conform to large-scale loading structures underlining the contact, which indicate still non-lithified shales when the overlying sandstones were deposited. Another possibility would be to suggest a flooding linked to a short-term, but significant regressive-transgressive interval, for instance to a glacio-eustatic event during the Early Ordovician. Going further, we note that the sequence of events linked to Unit 2 echoes the end-Ordovician development involving a glacio-isostatic flexure that allowed seawater to invade such a proximal area (see below). The Lower Ordovician inferred glaciers would have reached or at least been close to the Ennedi area. Although it required confirmation, this interpretation would explain well why the horizon of offshore shales characterizing the Unit 2 has been taken for a Lower Silurian interval by some previous geologists.

Whatever the age within the Ordovician, the proposed correlation framework depicts an overall south- to southeastward-oriented, continent-scale, progressive encroachment of the Lower Palaeozoic strata. It assumes the non-deposition and progressive disappearance of Middle-Upper Cambrian strata to the south and east, though correlative strata are well expressed to the north (northern Libya, north-Saharan basins of Algeria). The relatively monotonous coastal deposits preserved in the proximal reaches of the platform (Niger, SE Libya, Chad) are interpreted essentially as correlatives of transgressive Ordovician fine-grained inner-shelf to condensed successions in the north or NNW. Similar coastal suites are identified around Gondwana. In Jordan, they characterize an older upper Cambrian to Lower Ordovician interval [Disi Fm. Ghienne *et al.*, 2010, Meischner *et al.*, 2020], a time range logically expected in this intermediate, i.e., less proximal segment of the platform. In South Africa, the Peninsula Sandstone [Turner *et al.*, 2011] is assigned to a time interval comparable to that of the Lower Sandstones of the Ennedi. Interestingly, the Peninsula Sandstone characterizes a proximal setting where the succession wedge-out onto basement rocks at the scale

of a few tens of km, as should also occurred to the south of the Ennedi. Note that the relatively limited thickness (<200 m) of the Ordovician in the study area might not be solely a signature of the proximality of the Ennedi, as along-strike thickness changes reflecting subsidence/uplift trends controlled by basement heterogeneities [Nkono *et al.*, 2018] could also have played a significant role, as recognized elsewhere throughout the platform [Klitzsch, 1970, Perron *et al.*, 2021].

6.2. *The end-Ordovician glacial record*

The basal erosion surface and conglomerates of Unit 4, as well as the exhumed overlying glacial erosion surface and glaciomarine deposits at the base of Unit 5, allow both units to be interpreted as representatives of the end-Ordovician glacial record. Unit 4 (Wadi Djoua Fm.) appears as the diachronous(?) proximal counterpart of the Mamuniat Fm. of the Kufrah Basin [Le Heron *et al.*, 2010, Le Heron and Howard, 2012]. A “channelised unconformity” and an “intra-Mamuniat unconformity” in the Jebel Azbah area of the eastern Kufrah Basin might correlated with the basal bounding surfaces of Unit 4 and Unit 5, respectively.

The preservation of the glacial record is restricted to glacially-cut overdeepenings, possibly palaeovalleys, flanked by less eroded interfluvial areas, where Silurian–Devonian sandstones directly superimposed the preglacial Ordovician sandstones. The contrast between Unit 4 and Unit 5 is striking, the former dominated by conglomerate-prone outwash deposits, the latter being mainly a fine-grained subaqueous succession. Fining-up outwash deposits of Unit 4 suggest limited interferences with glacio-isostatic patterns during this earlier deglaciation phase, which is surprising in such a proximal setting. This might have been the case, however, if Unit 4 was principally deposited in the forefield of an ice sheet of a lesser extent (a less extensive glacial maximum? or a major stabilization? see discussion in Girard *et al.* [2015]). Fine-grained sedimentation of Unit 5 is more easily understandable as a consequence of a rapid ice-front retreat and deposition in a relatively and temporarily deep setting, both accounting for a major glacio-isostatic deflection in tens to hundreds of meters that inevitably interfered at that time with glacio-eustasy [Dietrich *et al.*, 2019]. The

fine-grained sandstone interval within Unit 5 might relate in one way or another to the late glacial recurrence identified on the eastern side of the Tibesti, which occurred before the deposition of Rhuddanian “warm” shales [Le Heron *et al.*, 2015, Meinhold *et al.*, 2016], which in turn can be correlated with the uppermost shale interval of Unit 5.

The temporal framework of the Bedo Fm. (Unit 5) has to be further discussed. The occurrence in the same stratigraphic position of an uppermost Ordovician to lowermost Silurian fauna at *c.* 500 km to the NNW [Le Hérisse *et al.*, 2013, Page *et al.*, 2013] and *c.* 650 km to the NNE [Thusu *et al.*, 2013] suggests the Unit 5 might be as old, or even slightly older—i.e., strictly latest Ordovician. The corresponding more distal deposits in the Kufrah Basin show well-developed more proximal HCS sandstone beds and a regressive evolution [Gindre *et al.*, 2012, Le Heron *et al.*, 2015] lacking in the Ennedi record. A latest Ordovician post-glacial, glacio-isostatically driven forced regression in the Ennedi area, with a coeval condensation to the north (or NW) might explain the distinct records. Such a setting predated the overall, long-term progradational trend characterizing the Lower Silurian [Bellini and Massa, 1980, Gindre *et al.*, 2012]. Therefore, the Bedo Fm. and Libyan analogues would be older than the Lower Silurian Tanezzuft shales usually defined in the Murzuq Basin [Bellini and Massa, 1980, Grignani *et al.*, 1991, Le Hérisse *et al.*, 2013, Lüning *et al.*, 1999], as are similar records preserved in southeastern Algeria and northern Niger [Denis *et al.*, 2010, Legrand, 2002].

7. Conclusion

The Lower Sandstones of the Ennedi is suggested to be essentially Ordovician in age. Here, the Palaeozoic succession illustrates the depositional conditions that occurred across one of the most proximal segments of the north-facing Gondwana platform. It shows in particular that the Ordovician transgressions, well-characterized in North Africa [Ghienne *et al.*, 2007], must have reached northern Chad. This archive also expands the area of glacially-controlled deposition to the south and east with respect to the pre-existing dataset. It confirms the eastward prolongation of the North African palaeoglacial framework in the eastern Sahara and, most likely, up to eastern Arabia and Ethiopia [Le Heron and Howard,

2012]. As far south as the Ennedi, glaciomarine sedimentation and northward-flowing ice flows are recognized. It indicates Ordovician ice-divide areas positioned in southern Chad or even further to the south in Central Africa. Here, a glacial record of unknown age but displaying northward-oriented dispersal patterns is identified [Censier and Lang, 1992], which might correlate with the Chadian Ordovician record. A potential end-Ordovician record have also been inferred in Cameroon, where an undated ice-flow pattern toward the WSW has been documented [Caron *et al.*, 2011]. Disparate orientations would not be a real problem because opposite flow orientations might characterize a same glacial event in areas corresponding to former ice divides [Rice *et al.*, 2020].

Conflicts of interest

Authors have no conflict of interest to declare.

Acknowledgments

The authors wish to thank the Governor of the Ennedi-Ouest and the Prefect of Fada, as well as the staff of the Ennedi African Park for their hospitality and for facilitating the fieldwork. They also gratefully acknowledge the University of N'Djaména, the CNRD (Centre National de Recherche pour le Développement), and the French Embassy in N'Djaména for their financial and technical support. The authors are grateful to G. Meinhold and J. N. Proust for their careful reviews, which helped improve an earlier version of the manuscript.

References

- Anfray, R. and Rubino, J.-L. (2003). Shelf depositional systems of the Ordovician Hawaz formation in the central Al Qarqaf high. In Salem, M. J., Oun, K. M., and Seddiq, H. M., editors, *The Geology of Northwest Libya II*, volume 3 of *Second Symposium on the Sedimentary Basin of Libya, Tripoli, Libya*, pages 123–134.
- Aubineau, J., El Albani, A., Chi Fru, E., Gingras, M., Batonneau, Buatois, L. A., Geffroy, C., Labanowski, J., Laforest, C., Lemée, L., Mangano, M. G., Meunier, A., Pierson-Wickmann, A.-C., Recourt, P., Riboulleau, A., Trentesaux, A., and Konhauser, K. O. (2018). Unusual microbial mat-related structural diversity 2.1 billion years ago and implications for the Francevillian biota. *Geobiology*, 16, 476–497.
- Bataller, F. J., McDougall, N., and Moscariello, A. (2019). Ordovician glacial paleogeography: Integration of seismic spectral decomposition, well sedimentological data, and glacial modern analogs in the Murzuq Basin, Libya. *Interpretation*, 7, T383–T408.
- Bellini, E., Giori, I., Ashuri, O., and Benelli, F. (1991). Geology of Al Kufrah Basin, Libya. In Salem, M. J. and Belaid, M. N., editors, *The Geology of Libya*, volume V, pages 2155–2184. Academic Press, London.
- Bellini, E. and Massa, D. (1980). A stratigraphic contribution to the Palaeozoic of Southern basins of Libya. In Salem, M. J. and Buswiel, M. T., editors, *The Geology of Libya*, volume I, pages 3–56. Academic Press, London.
- Beuf, S., Biju-Duval, B., De Charpal, O., Rognon, P., Gariel, O., and Bennacef, A. (1971). Les grès du Paléozoïque inférieur au Sahara. In *Science et Technique du Pétrole*, volume 18, page 465. Ed. Technip, Paris.
- Bizard, C., Bonnet, A., Freulon, J.-M., Gerard, G., De Lapparent, A., Vincent, P., and Wacrenier, P. (1955). La série géologique entre le Djado et le Tibesti (Sahara oriental). *C. R. Acad. Sci. Paris*, 241, 1320–1323.
- Bonnet, A., Freulon, J.-M., de Lapparent, A.-F., and Vincent, P. (1955). Observations géologiques sur l'ennedi, le Mourdi et les Erdi (Territoire du Tchad, A.E.F.). *C. R. Acad. Sci. Paris*, 241, 1403–1405.
- BRGM (2004). *Carte offerte à l'occasion du 20^{ème} Colloque de Géologie Africaine (Orléans)*. BRGM Editions, Orléans.
- Buatois, L. A. and Mangano, M. G. (2011). *Ichnology: Organism-Substrate Interactions in Space and Time*. Cambridge University Press, Cambridge.
- Caron, V., Mahieux, G., Ekomane, E., Moussango, P., and Babinski, M. (2011). One, two or no record of late neoproterozoic glaciation in South-East Cameroon. *J. Afr. Earth Sci.*, 59, 111–124.
- Censier, C. and Lang, J. (1992). La Formation glaciaire de la Mambéré (République Centrafricaine): Reconstitution paléogéographique et implications à l'échelle du Paléozoïque africain. *Geol. Rund.*, 81, 769–789.
- Dabard, M. P., Loi, A., Paris, F., Ghienne, J.-F., Pis-

- tis, M., and Vidal, M. (2015). Sea-level curve for the Middle to early Late Ordovician in the Armorican Massif (western France): Icehouse third-order glacio-eustatic cycles. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 436, 96–111.
- Dashtgard, S. E., Gingras, M. K., and MacEachern, J. A. (2009). Tidally modulated shorefaces. *J. Sediment. Res.*, 79, 793–807.
- De Lestang, J. (1965). Das Paläozoikum am Rande des Afro-Arabischen Gondwanakontinents. *Z. Deutsch. Geol. Ges.*, 117, 479–488.
- Denaeyer, M.-E. (1924). L'Ouadaï oriental et les régions voisines. Géographie physique, géologie, lithologie, d'après les documents de la mission de délimitation Ouadaï-Darfour. *Bull. Soc. Géol. France*, 24, 538–576.
- Denis, M., Guiraud, M., Konaté, M., and Buoncristiani, J.-F. (2010). Subglacial deformation and water-pressure cycles as a key for understanding ice stream dynamics: evidence from the Late Ordovician succession of the Djado Basin (Niger). *Intern. J. Earth Sci.*, 99, 1399–1425.
- Deschamps, R., Eschard, R., and Roussé, S. (2013). Architecture of Late Ordovician glacial valleys in the Tassili N'Ajjer area (Algeria). *Sediment. Geol.*, 289, 124–147.
- Deynoux, M., Sougy, J., and Trompette, R. (1985). Lower Palaeozoic rocks of west Africa and the western part of central Africa. In Holland, C. H., editor, *Lower Palaeozoic of North-Western and West-Central Africa*, pages 337–495. John Wiley & Sons, Chichester.
- Dietrich, P., Ghienne, J. F., Lajeunesse, P., Normandeau, A., Deschamps, R., and Razin, P. (2019). Deglacial sequences and glacio-isostatic adjustment: Quaternary compared with Ordovician glaciations. *Geol. Soc. Lond. Spec. Publ.*, 475, 149–179.
- Elicki, O. and Altumi, M. M. (2019). Cambrian trace fossils from North Africa and their contribution to Gondwana's paleobiogeography and depositional history. *J. Afr. Earth Sci.*, 158, article no. 103556.
- Eschard, R., Braik, F., Bekkouche, D., Rahuma, M. B., Desaubliaux, G., Deschamps, R., and Proust, J. N. (2010). Palaeohighs: their influence on the North African Palaeozoic petroleum systems. In *Petroleum Geology Conference Series*, volume 7, pages 707–724. Geological Society, London.
- Eyles, N., Putkinen, N., Sookhan, S., and Arbelaez-Moreno, L. (2016). Erosional origin of drumlins and megaridges. *Sediment. Geol.*, 338, 2–23.
- Fabre, J. and Kazi-Tani, N. (2005). Ordovicien, Silurien, Devonien, Permo-Carbonifère. In Fabre, J., editor, *Géologie du Sahara Occidental et Central, Tervuren African Geoscience Collection*, volume 18, pages 147–360. Musée Royal de l'Afrique Centrale, Tervuren, Belgique.
- Fritel, P.-H. (1924). Sur les restes de végétaux fossiles paléozoïques recueillis en Ouadaï par la mission du L^t-Col. Grossard. *Bull. Mus. Hist. Nat.*, 1, 117–118.
- Fritel, P.-H. (1925). Végétaux paléozoïques et organismes problématiques de l'Ouadaï. *Bull. Soc. Géol. France*, 25, 33–48.
- Gatinskiy, G., Klochko, V. P., Rozman, K. S., and Trofimov, D. M. (1966). Nouvelles données sur la stratigraphie des dépôts paléozoïques du Sahara méridional. *Dokl. Akad. Nauk SSSR*, 170, 1154–1157.
- Germann, K., Wipki, M., and Schwarz, T. (1993). Cambro-Ordovician bauxitic laterites of NW-Sudan. In Thorweih, U. and Schandemeier, H., editors, *Geoscientific Research in Northeastern Africa*, pages 335–340. Balkema, Amsterdam.
- Ghienne, J.-F., Boumendjel, K., Paris, F., Videt, B., Racheboeuf, P., and Ait Salem, H. (2007). The Cambrian-Ordovician succession in the Ougarta Range (western Algeria, North Africa) and interference of the Late Ordovician glaciation on the development of the Lower Palaeozoic transgression on northern Gondwana. *Bull. Geosci.*, 82(3), 183–214.
- Ghienne, J.-F., Monod, O., Kozlu, H., and Dean, W. T. (2010). Cambrian-Ordovician depositional sequences in the Middle East: a perspective from Turkey. *Earth Sci. Rev.*, 101, 101–146.
- Ghienne, J.-F., Moreau, J., Degermann, L., and Rubino, J.-L. (2013). Lower Palaeozoic unconformities in an intracratonic platform setting: Glacial erosion vs. tectonics in the eastern Murzuq Basin (southern Libya). *Int. J. Earth Sci.*, 102, 455–482.
- Gil-Ortiz, M., McDougall, N. D., Cabello, P., Marzo, M., and Ramos, E. (2022). Sedimentary architecture of a Middle Ordovician embayment in the Murzuq Basin (Libya). *Mar. Pet. Geol.*, 135, article no. 105339.
- Gindre, L., Le Heron, D., and Bjørnseth, H. M. (2012). High resolution facies analysis and sequence stratigraphy of the Siluro-Devonian succession of Al Kufrah basin (SE Libya). *J. Afr. Earth*

- Sci.*, 76, 8–26.
- Girard, F., Ghienne, J.-F., Du-Bernard, X., and Rubino, J.-L. (2015). Sedimentary imprints of former ice-sheet margins: Insights from an end-Ordovician archive (SW Libya). *Earth Sci. Rev.*, 148, 259–289.
- Greigert, J. and Pougnet, R. (1967). Essai de description des formations géologiques de la République du Niger. In *Direction des Mines et de la Géologie 3 (République du Niger)*. BRGM Ed, Paris, France.
- Grignani, D., Lanzoni, E., and Elatrash, H. (1991). Palaeozoic and Mesozoic subsurface palynostratigraphy in the Al Kufrah Basin, Libya. In Salem, M. J., Hammuda, O. S., and Eliagoubi, B. A., editors, *The Geology of Libya*, volume IV, pages 1160–1227. Elsevier, Amsterdam.
- Hart, B. S. and Plint, A. G. (1995). Gravelly shoreface and beachface deposits. *Spec. Publ. Int. Ass. Sediment.*, 22, 75–99.
- Jouliia, F. (1959). Les series primaires au N et NW de l’Aïr (Sahara central). Discordances observées. *Bull. Soc. Géol. France*, 7, 192–196.
- Kichou-Braïk, F., Samar, L., Fekirine, B., and Legrand, P. (2006). Découverte de graptolites d’âge caradozien dans quelques sondages du Tinrhert (Sahara algérien). *C. R. Palevol*, 5, 675–683.
- Kilian, C. (1924). Notes sur la géologie du Sahara central. *Trav. Lab. Géol. Fac. Sc. Univ. Grenoble*, 13, 87–102.
- Klitzsch, E. (1965). Comments on the Geology of the central parts of Southern Libya and Northern Chad. *Pet. Explor. Soc. Libya*, 8, 1–17.
- Klitzsch, E. (1968). Die Gotlandium-Transgression in der Zentral-Sahara. *Z. Deutsch. Geol. Ges.*, 117, 492–501.
- Klitzsch, E. (1970). Die Strukturgeschichte der Zentralsahara. *Geol. Rund.*, 59, 459–527.
- Klitzsch, E. (1994). Geological exploration history of the Eastern Sahara. *Geol. Rund.*, 83, 475–483.
- Klitzsch, E., Reynolds, P.-O., and Barazi, N. (1993). Géologique Erkundungsfahrt zum Ostteil des Ennedi Gebirges und der Mourdiu Depression (NE Sudan) im Januar 1992. *Würzbg. Geogr. Arb.*, 87, 49–61.
- Lacroix, A. and Tilho, J. (1919). Esquisse géologique du Tibesti, du Borkou, de l’Erdi et de l’Ennedi—Les formations sédimentaires. *C. R. Acad. Sci., Paris*, 168, 1169–1174.
- Lang, J., Le Heron, D. P., Van den Berg, J. H., and Winsemann, J. (2021). Bedforms and sedimentary structures related to supercritical flows in glacial settings. *Sedimentology*, 68, 1539–1579.
- Le Hérisse, A., Paris, F., and Steemans, P. (2013). Late Ordovician-earliest Silurian palynomorphs from northern Chad and correlation with contemporaneous deposits of southeastern Libya. *Bull. Geosci.*, 83, 483–504.
- Le Heron, D. P. (2018). An exhumed Paleozoic glacial landscape in Chad. *Geology*, 46, 91–94.
- Le Heron, D. P., Armstrong, H. A., Wilson, C., Howard, J. P., and Gindre, L. (2010). Glaciation and deglaciation of the Libyan Desert: The Late Ordovician record. *Sediment. Geol.*, 223, 100–125.
- Le Heron, D. P., Busfield, M. E., Chen, X., Corkeron, M., Davies, B. J., Dietrich, P., Ghienne, J.-F., Kettler, C., Scharfenberg, L., Vandyk, T. M., and Wohlschägl, R. (2022). New perspectives on glacial geomorphology in Earth’s deep time record. *Front. Earth Sci.*, 10, article no. 870359.
- Le Heron, D. P. and Howard, J. P. (2012). Sandstones, glaciers, burrows and transgressions: the lower Palaeozoic of Jabel az-Zalmah, Al Kufrah basin, Libya. *Sediment. Geol.*, 245, 63–75.
- Le Heron, D. P., Meinhold, G., Elgadry, M., Abutaruma, Y., and Boote, D. (2015). Early Palaeozoic evolution of Libya: perspectives from Jabal Eghei with implications for hydrocarbon exploration in Al Kufrah Basin. *Basin Res.*, 27, 60–83.
- Le Heron, D. P., Sutcliffe, O. E., Whittington, R. J., and Craig, J. (2005). The origins of glacially related soft-sediment deformation structures in Upper Ordovician glaciogenic rocks: implication for ice-sheet dynamics. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 218, 75–103.
- Leckie, D. (1988). Wave-formed, coarse-grained ripples and their relationship to hummocky cross-stratification. *J. Sediment. Res.*, 58, 607–622.
- Legrand, P. (1974). Essai sur la paléogéographie de l’Ordovicien au Sahara algérien. *Notes Mém. Cie. fr. pét.*, 11, 121–138.
- Legrand, P. (2002). Bâtir une stratigraphie: les leçons de l’étude du Paléozoïque au Sahara algérien. *C. R. Palevol*, 1, 383–397.
- Lemoigne, Y., Durante, M., and Blanc, C. (1992). Flores d’âge Carbonifère Inférieur du Niger et du Tchad. *Geobios*, 25, 449–455.
- Lessertisseur, J. (1956). Sur un bilobite nouveau du Gothlandien de l’Ennedi (Tchad, A.E.F.). *Cruziana ancora*. *Bull. Soc. Géol. France*, 6, 43–47.
- Liégeois, J.-P., Abdelsalam, M. G., Ennih, N., and

- Ouabadi, A. (2013). Metacraton: nature, genesis and behavior. *Gondwana Res.*, 23, 220–237.
- Lucchi, R. G., Sagnotti, L., Camerlenghi, A., Macri, P., Rebesco, M., Pedrosa, M. T., and Giorgetti, G. (2015). Marine sedimentary record of Meltwater Pulse 1a along the NW Barents Sea continental margin. *Arktos*, 1, article no. 7.
- Lüning, S., Craig, J., Fitches, B., Mayouf, J., Busrewil, A., El Dieb, M., Gammudi, A., Loydell, D., and McIlroy, D. (1999). Re-evaluation of the petroleum potential of the Kufra Basin (SE Libya, NE Chad): does the source rock barrier fall? *Mar. Pet. Geol.*, 16, 693–718.
- Lüning, S., Miles, N., Pearce, T., Brooker, E., Barnard, P., Johannson, G., and Schäfer, S. (2010). Biostratigraphy, chemostratigraphy and thermal maturity of the A1-NC198 exploration well in the Kufra Basin, SE Libya. In *Petroleum Geology Conference Series*, volume 7, pages 761–770. Geology Society of London, London.
- MEH (Ministère de l'Élevage et de l'Hydraulique) (2015). Carte hydrogéologique de la République du Tchad au 1:200 000, Ouvrages et Ressources, feuille NE-34-10 Fada. UNITAR and SWISSTOPO, Genève and Wabern.
- Meinhold, G., Le Heron, D. P., Elgadry, M., and Abutarruma, Y. (2016). The search for 'hot shales' in the western Kufra Basin, Libya: geochemical and mineralogical characterisation of outcrops, and insights into latest Ordovician climate. *Arab. J. Geosci.*, 9, article no. 62.
- Meischner, T., Elicki, O., Masri, A., Moumani, K. A., and Hussein, M. A. A. (2020). Ordovician trace fossils from southern Jordan with particular consideration to the Cruziana rugosa group: Taxonomy, stratigraphy and trans-regional correlation throughout the Middle East and northern Africa. *J. Afr. Earth Sci.*, 164, article no. 103595.
- Moreau, J., Ghienne, J.-F., Le Heron, D., Rubino, J.-L., and Deynoux, M. (2005). A 440 Ma old ice stream in North Africa. *Geology*, 33, 753–756.
- Newton, A. M. and Huuse, M. (2017). Glacial geomorphology of the central Barents Sea: Implications for the dynamic deglaciation of the Barents Sea Ice Sheet. *Mar. Geol.*, 387, 114–131.
- Nkono, C., Liégeois, J. P., and Demaiffe, D. (2018). Relationships between structural lineaments and Cenozoic volcanism, Tibesti swell, Saharan metacraton. *J. Afr. Earth Sci.*, 145, 274–283.
- Page, A., Meinhold, G., Le Heron, D. P., and Elgadry, M. (2013). *Normalograptus kufraensis*, a new species of graptolite from the western margin of the Kufra Basin, Libya. *Geol. Mag.*, 150, 743–755.
- Perron, P., Guiraud, M., Vennin, E., Moretti, I., Portier, E., Le Pourhiet, L., and Konaté, M. (2018). Influence of basement heterogeneity on the architecture of low subsidence rate Paleozoic intracratonic basins (Reggane, Ahnet, Mouydir and Illizi basins, Hoggar Massif). *Solid Earth*, 9, 1239–1275.
- Perron, P., Le Pourhiet, L., Guiraud, M., Vennin, E., Moretti, I., Portier, É., and Konaté, M. (2021). Control of inherited accreted lithospheric heterogeneity on the architecture and the low, long-term subsidence rate of intracratonic basins. *BSGF-Earth Sci. Bull.*, 192, article no. 15.
- Proust, J. N., Pouderoux, H., Ando, H., Hesselbo, S. P., Hodgson, D. M., Lofi, J., Rabineau, M., and Sugarman, P. J. (2018). Facies architecture of Miocene subaqueous clinothems of the New Jersey passive margin: Results from IODP-ICDP Expedition 313. *Geosphere*, 14, 1564–1591.
- Ravier, E. and Buoncristiani, J. F. (2018). Glaciohydrogeology. In *Past Glacial Environments*, pages 431–466. Elsevier Ltd, Amsterdam, 2nd edition.
- Rice, J. M., Ross, M., Paulen, R. C., Kelley, S. E., and Briner, J. P. (2020). A GIS-based multi-proxy analysis of the evolution of subglacial dynamics of the Quebec–Labrador ice dome, northeastern Quebec, Canada. *Earth Surf. Proc. Land.*, 45, 3155–3177.
- Riché, P. and de Charpal, O. (1998). An example of platform sedimentation: the Saharan Cambro-Ordovician. In Majithia, M., editor, *Dynamics and Methods of Study of Sedimentary Basins, Editions Technip—Association des sédimentologues français*, pages 91–110. Ed. Technip, Paris. translation of *Dynamique et méthodes d'étude des bassins sédimentaires*, 1989.
- Rubino, J.-L., Anfray, R., Blanpied, C., Ghienne, J.-F., and Manatschal, G. (2003). Meander belt complex within the lower Mamuniyat Formation in western Al Qarqaf area, Libya. In Salem, M. J. and Khaled Oun, M., editors, *The Geology of Northwest Libya*, volume II, pages 3–18. Gutenberg Press, Malta.
- Rüther, D. C., Bjarnadóttir, L. R., Junntila, J., Husum, K., Rasmussen, T. L., Lucchi, R. G., and Andreassen, K. (2012). Pattern and timing of the northwestern Barents Sea Ice Sheet deglaciation and indications

- of episodic Holocene deposition. *Boreas*, 4, 494–512.
- Sandford, K. S. (1935). Geological Observations on the Northwest Frontiers of the Anglo-Egyptian Sudan and the adjoining part of the Southern Libyan Desert. *Q. J. Geol. Soc.*, 91, 323–381.
- Schneider, J.-L. and Wolff, J.-P. (1992). Carte Géologique et Carte Hydrogéologique 1/1500000 de la République du Tchad, Mémoire Explicatif. In *Documents du BRGM*, volume 209, page 531. Edition BRGM, Orléans, France.
- Seilacher, A. (1970). Cruziana stratigraphy of ‘non-fossiliferous’ Paleozoic sandstones. In Crimes, T. P. and Harper, J. W., editors, *Trace Fossils*, volume 3 of *Geological Journal Special Issue*, pages 447–476.
- Seilacher, A. (2007). *Trace Fossil Analysis*. Springer, Berlin.
- Seilacher, A., Lüning, S., Martin, M. A., Klitzsch, E., Khoja, A., and Craig, J. (2002). Ichnostratigraphic correlation of lower Palaeozoic clastics in the Kufra Basin (SE Libya). *Lethaia*, 35, 257–262.
- Sharland, P. R., Casey, D. M., Davies, R. B., Simmons, M. D., and Sutcliffe, O. E. (2001). *Arabian Plate Sequence Stratigraphy (GeoArabia Spec. Publication 2)*. Gulf PetroLink, Bahrain.
- Smith, J. A., Graham, A. G., Post, A. L., Hillenbrand, C. D., Bart, P. J., and Powell, R. D. (2019). The marine geological imprint of Antarctic ice shelves. *Nat. Commun.*, 10, article no. 5635.
- Spagnolo, M., Clark, C. D., Ely, J. C., Stokes, C. R., Anderson, Andreassen, K., Graham, A. G. C., and King, E. C. (2014). Size, shape and spatial arrangement of mega-scale glacial lineations from a large and diverse dataset. *Earth Surf. Process. Landf.*, 39, 1432–1448.
- Thomas, K., Herminghaus, S., Porada, H., and Goehring, L. (2013). Formation of *Kinneyia* via shear-induced instabilities in microbial mats. *Philos. Trans. R. Soc. A*, 371, article no. 20120362.
- Thusu, B., Rasul, S., Paris, F., Meinhold, G., Howard, J. P., Abutarruma, Y., and Whitham, A. G. (2013). Latest Ordovician–earliest Silurian acritarchs and chitinozoans from subsurface samples in Jebel Asba, Kufra Basin, SE Libya. *Rev. Palaeobot. Palynol.*, 197, 90–118.
- Trompette, R. (1983). Le Paléozoïque du Niger et du Tchad. In Fabre, J., editor, *West Africa, Geological Introduction and Stratigraphic Terms*, pages 121–126. Pergamon Press, Oxford.
- Turner, B. R. (1980). Palaeozoic sedimentology of the southeastern part of Al Kufrah Basin, Libya: a model for oil exploration. In Salem, M. J. and Busrewil, M. T., editors, *The Geology of Libya*, volume II, pages 351–374. Academic Press, London.
- Turner, B. R. (1991). Palaeozoic deltaic sedimentation in the southeastern part of Al Kufrah Basin, Libya. In Salem, M. J. and Belaid, M. N., editors, *The Geology of Libya*, volume V, pages 1713–1726. Elsevier, Amsterdam.
- Turner, B. R., Armstrong, H. A., and Holt, P. (2011). Visions of ice sheets in the early Ordovician greenhouse world: Evidence from the Peninsula Formation, Cape Peninsula, South Africa. *Sediment. Geol.*, 236, 226–238.
- Wacrenier, P. (1958). *Notice explicative de la carte géologique provisoire du Borkou-Ennedi-Tibesti au 1 : 1 000 000*. Dir. des Mines et de la Géologie (AEF), Brazzaville, Congo.
- Wolff, J.-P. (1964). *Carte Géologique de la République du Tchad*. BRGM, Paris, France.
- Zecchin, M., Catuneanu, O., and Caffau, M. (2007). High-resolution sequence stratigraphy of clastic shelves V: criteria to discriminate between stratigraphic sequences and sedimentological cycles. *Mar. Pet. Geol.*, 85, 259–271.