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Fault slip of the 2022 Mw6.7 Menyuan, China earthquake observed by InSAR, and its tectonic implications

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In this article the ascending and descending Sentinel-1A satellite data are used to investigate the coseismic slip model of the 2022 Mw6.7 Menyuan earthquake in Qinghai, China. The optimal slip model indicates that this event ruptured two fault segments. The main rupture concentrated on the western Lenglongling fault (LLLF) with a purely left-lateral striking-slip motion. A small part of the eastern Tuolaishan fault (TLSF) section was also ruptured, and the motion on it is mainly oblique slip at depth, with an obvious thrust component. Combined with the rupture characteristics of historical events, GPS velocity map, and slip rate studies, we suggest that the TLSF–LLLF junction is a demarcation point where the deformation partitioning pattern has changed. Along the whole LLLF segment, the oblique convergence has completely partitioned into slip on the purely strike-slipping LLLF and thrusting faults in the north. The TLSF segment accommodates a fraction of compressional shortening, which compensates for the discrepancy in the left-lateral slip rate between the LLLF and TLSF. Such transformation in the strain partitioning pattern is likely to be determined by the geometric relationship between the fault strike and the direction of regional block movement.

KEYWORDS

the Lenglongling fault, the Tuolaishan fault, the 2022 Mw6.7 Menyuan earthquake, Menyuan earthquake, InSAR, deformation partitioning pattern

Introduction

On 8 January 2022, an Mw6.7 earthquake struck ~59 km far away from Menyuan County in the Qinghai Province of northwestern China at the western section of the Qilian–Haiyuan fault (QL-HYF) zone. According to the focal-mechanism solution reported by the United States Geological Survey (USGS) and Global Centroid Moment Tensor (GCMT), the fault movement caused by the main shock was dominated by an NWW left-lateral strike-slip with a thrust component (Table 1; Figure 1).

As the leading edge of the northeastern Tibetan Plateau, the QL-HYF zone accommodates about a quarter of the convergence rate driven by the ongoing collision of the India and Eurasia plates, as shown in previous investigations (e.g.,

TABLE 1 Focal mechanisms of the 2022 Menyuan earthquake reported by different studies.

Source	Location of the epicenter		Depth (km)	Focal mechanism		Mag. (<i>M_w</i>)
	Lon. (°E)	Lat. (°N)		(Strike/dip/rake) (°)		
USGS	101.29	37.83	13.0	NP 1	13°/75°/178°	6.61
				NP 2	104°/88°/15°	
GCMT	101.31	37.80	14.8	NP 1	14°/89°/172°	6.70
				NP 2	104°/82°/1°	
This study	101.27	37.80	~3.0	NP 1	-	6.69
				NP 2	115°/88°/1°	

USGS: United States Geological Survey; GCMT: Global Centroid Moment Tensor.

Molnar & Tapponnier, 1975; England & Houseman, 1985; Tapponnier et al., 1990; Pichon et al., 1992; England & Molnar, 1997; Tapponnier et al., 2001; Flesch et al., 2001; Yuan et al., 2013; Zuza et al., 2018). The strain distribution of the NE Tibetan Plateau from the recent Global Navigation Satellite System (GNSS) and interferometric synthetic aperture radar (InSAR) observations (Wang and Shen, 2020; Huang et al., 2022; Ou et al., 2022; Figure 1B) shows dominant shearing deformation on the middle of the northeast edge of the Tibetan Plateau and extrusion deformation on the eastern and western areas. The left-lateral shearing deformation largely focuses on the QL–HYF zone. Previous studies have suggested that an ESE trending continuously transferred deformation along fault segments of the QL–HYF zone from west to east (Gaudermer et al., 1995; Lasserre et al., 2002; Yuan et al., 2004; Zheng et al., 2013; Yao et al., 2019). However, recent studies do not seem to support such a tectonic transformation model. Field geological investigations at the middle QL–HYF zone show that the strain on the middle Lenglongling fault (LLLF) tends to be transferred northeastward to the eastern LLLF, rather than transmitted eastward to the Jinqianghe fault (JQHF) (Guo et al., 2019, 2020; Gao et al., 2021). InSAR observations from Envisat and Sentinel-1 satellites present slip partitioning and segmented movement for the 1,000-km QL–HYF system (Daout et al., 2016; Huang et al., 2022; Ou et al., 2022). As for the western section of the QL–HYF (the TLSF and further west), the fault slip rate and strain partitioning have not been properly investigated. Previous research (IGLIS, 1993; Yuan et al., 2003; Deng et al., 2007; Xu et al., 2016) concentrated on the left-lateral shear movement on the TLSF and suggested a dominant sinistral strike-slip on the TLSF. Huang et al. (2022) suggested that the Qilianshan absorbs ~4 mm/yr left-lateral shear discrepancy between the TLSF and middle segment of the QL–HYF by shortening, given a low dip-slip rate on the TLSF from the InSAR profile inversion. However, the recent cross-fault GPS and leveling profiles show an obvious shortening and vertical deformation across the TLSF (Li et al., 2022, Supplementary Figure S1). It is controversial how the strain is

partitioned along the western section of the QL–HYF and other faults in the north. The present-day slip rates estimated from InSAR cross-fault profiles (Huang et al., 2022; Ou et al., 2022) show an obvious step between the Tuolaishan fault (TLSF) segment and LLLF segment, where the 2022 Menyuan event occurred. It is a good opportunity to investigate the tectonic role of the TLSF and LLLF sections in the deformation transformation model of the northeast edge of the Tibetan Plateau from this earthquake.

In this study, both the ascending and descending interferograms from the Sentinel-1A satellite, whose tropospheric delays were corrected by a method dedicated to the small-to-moderate-magnitude earthquake proposed by Gong et al. (2022), were used to investigate the coseismic displacements of the Menyuan earthquake. Combined with previous geodetic and geological observations, we try to 1) determine the slip kinematics of the 2022 Mw6.7 Menyuan earthquake and discuss its tectonic implications and 2) analyze how the tectonic deformation is transformed from the TLSF to LLLF. Our results will help to improve understanding of the deformation partitioning patterns in the western section of the northeastern edge of the Tibetan Plateau.

Tectonic setting

As a major tectonic structure in the northeastern margin of the Tibetan Plateau, the QL–HYF zone plays an important role in accommodating the northeastward expansion of the Tibetan Plateau into the continental interior (e.g., Lasserre et al., 2002; Duvall and Clark, 2010; Zheng et al., 2013; Guo et al., 2017; Huang et al., 2022). Located at the convergence boundaries between Tibet, Gobi–Alashan block and Ordos block, it presents a geometrical complexity along faults, composed of several segments (Figures 1A,B), i.e., the TLSF, LLLF, JQHF, Maomaoshan fault (MMSF), Laohushan fault (LHSF), Haiyuan fault (HYF), and Liupanshan fault (LPSF) from west to east. Fault

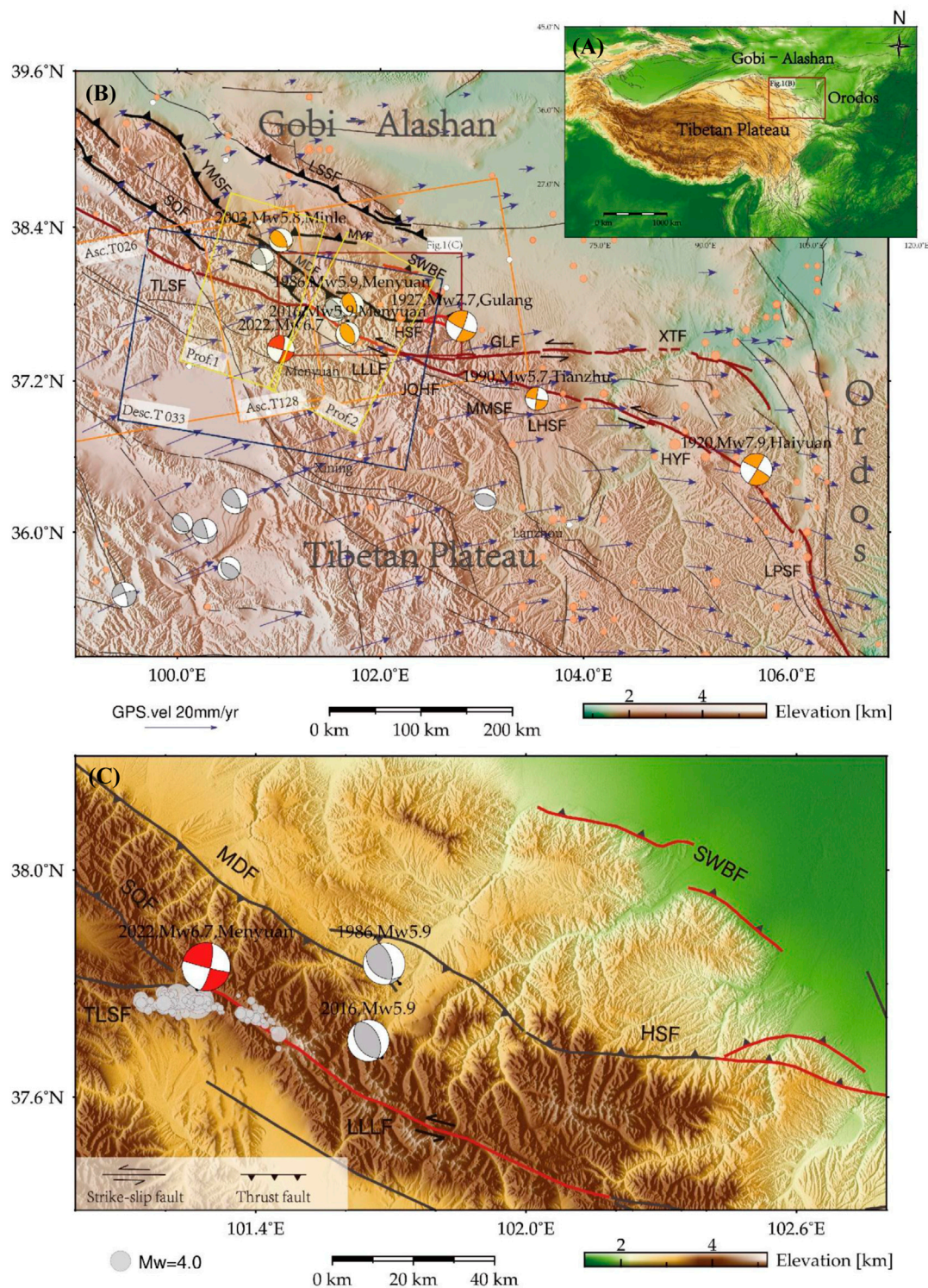


FIGURE 1
 Tectonic setting of the 2022 Mw6.7 Menyuan earthquake. (A) Tectonic setting of the Qilian-Haiyan fault (QL-HYF) zone. (B) Main faults in the NE Tibetan Plateau. Active faults are from Deng et al. (2007) (<https://www.activefault-datacenter.cn/map>). The Mw 5.7 + historical earthquakes that occurred along the QL-HYF zone since 1920 are marked by the GCMT focal mechanisms. Arrows indicate GPS velocities from Wang and Shen, (2020). The orange and blue rectangles outline the coverage of ascending and descending Sentinel-1 SAR data. Two yellow rectangles outline the coverage of the GPS profiles across the TLSF and LLLF, which are shown in Supplementary Figure S1. TLSF: Tuolaishan fault; LLLF: Lenglongling fault; JQHF: Jinqianghe fault; MMSE: Maomaoshan fault; LHSF: Laohushan fault; HYF: Haiyuan fault; LPSF: Liupanshan fault; GLF: Gulang fault; XTF: (Continued)

FIGURE 1 (Continued)

Xiangshan–Tianjingshan fault; HSF: Huangcheng–Shuangta fault; SWBF: Southern Wuwei Basin fault; SQF: Sunan–Qilian fault; YMSF: Yumushan fault; MDF: Minle–Damaying fault; MYF: Minle–Yongchang fault; LSSF: Longshoushan fault. **(C)** Tectonic map in the seismogenic area of the 2022 Mw6.7 Menyuan earthquake. The surface rupture of the 1927 Gulang earthquake in red is modified from Guo et al. (2020). The gray circles present the aftershocks within the first 10 days after the main shock was recorded by the local seismometer array (Fan et al., 2022). Two moderate-thrust earthquakes that occurred in this region in 1986 and 2016 are marked by the GCMT focal mechanisms.

TABLE 2 Focal mechanisms of the historical tectonic events ($M_w > 5.7$) in the QL–HYF zone.

Tectonic event	Location of the epicenter		Depth (km)	Focal mechanism (Strike/dip/rake) (°)	Mag. (M_w)	Rupture type
	Lon. (°E)	Lat. (°N)				
1920, Haiyuan	105.54	36.48	6.0	110°/90°/10°	7.9	Strike-slip
1927, Gulang	102.37	38.05	10.0	110°/45°/45°	7.7	Strike-slip with slight thrust
1986, Menyuan	101.72	37.80	15.0	346°/60°/113°	5.9	Thrust
1990, Tianzhu	10.354	37.06	15.0	98°/85°/–3°	5.7	Strike-slip
2003, Minle	101.02	38.30	15.0	331°/58°/116°	5.8	Thrust with slight strike-slip
2016, Menyuan	101.68	37.67	14.3	335°/47°/96°	5.9	Thrust

The focal mechanism of the 1920 Haiyuan earthquake is sourced from Ou et al. (2020), and the focal mechanism of the 1927 Gulang earthquake is sourced from Molnar and Deng (1984). Others are sourced by Global Centroid Moment Tensor.

kinematics and seismicity on those fault segments have been extensively studied by using both geological and geodetic observations over the last decades. Although a high west-to-east decreasing strike-slip rate (19 ± 5 mm/yr and 12 ± 4 mm/yr) is suggested for the QL–HYF by early geological studies (Zhang et al., 1988; Lasserre et al., 1999; Lasserre et al., 2002), the follow-on research studies give a relatively consistent estimate with an arc-shaped distribution, from 0–2 mm/yr in the western end to 4.5–9 mm/yr in the middle section, and then decreasing to 0–2 mm/yr in the eastern end near the LPSF (e.g., He et al., 2001; Gan et al., 2007; Li et al., 2009; Duvall and Clark, 2010; Loveless and Meade, 2011; Zheng et al., 2013; Jiang et al., 2017; Liu et al., 2018; Shao et al., 2020). InSAR observations (Daout et al., 2016; Song et al., 2019; Huang et al., 2022; Ou et al., 2022) provide a detailed slip rate and slip partitioning map on the QL–HYF and some northerly located faults, such as the Gulang fault (GLF), Xiangshan–Tianjingshan fault (XTF), Huangcheng–Shuangta fault (HSF), Southern Wuwei Basin fault (SWBF), Sunan–Qilian fault (SQF), Yumushan fault (YMSF), Minle–Damaying fault (MDF), Minle–Yongchang fault (MYF), and Longshoushan fault (LSSF) (Figures 1B,C). The QL–HYF zone was seismically active in the past century. More than six $M_w 5.5+$ earthquakes occurred along it, including the 1920 $M_w 7.9$ Haiyuan earthquake, 1927 $M_w 7.7$ Gulang earthquake, 1986 $M_w 5.9$ Menyuan earthquake,

1990 $M_w 5.7$ Tianzhu earthquake, 2003 $M_w 5.8$ Minle earthquake, and 2016 $M_w 5.9$ Menyuan earthquake (Table 2; Figures 1B,C).

The seismogenic fault of the 2022 $M_w 6.7$ Menyuan earthquake is the TLSF–LLLF, the western section of the QL–HYF. Tectonically, the LLLF branches into the TLSF and northerly located SQF at its western end, and the TLSF steps left in a ~3-km left-stepped en echelon pattern (Guo et al., 2019; Pan et al., 2022). The TLSF and further west have not been properly investigated due to inaccessibility and remoteness, and it is suggested to be dominated by a sinistral strike-slip (Deng et al., 2007). The recent InSAR results (Huang et al., 2022) gave a strike-slip rate map along the TLSF, and its along-strike-slip rate gradually increases from 1.8 ± 0.3 mm/yr close to Halahu in the western end, to 2.8 mm/yr in the middle, and 3.5 mm/yr along the easternmost segment. The motion on the LLLF is predominantly left-lateral strike-slip during the late Quaternary (Guo et al., 2020), following a thrusting movement due to the northeastward compression in the early Quaternary (He et al., 2001; Li et al., 2009). Most studies of the geological slip rate of the LLLF focus on the middle section. Among those studies, the two most recent studies from Jiang et al. (2017) and Guo et al. (2017) reported a consistent estimate with a slip rate of 6.6 ± 0.3 mm/yr and 6.4 ± 0.7 mm/yr, respectively, in agreement with InSAR-inverted one (6.4 ± 0.5 mm/yr) by Huang et al. (2022). Gao et al. (2021) gave a slip rate of 6.0 ± 0.8 mm/yr for the eastern section of the LLLF,

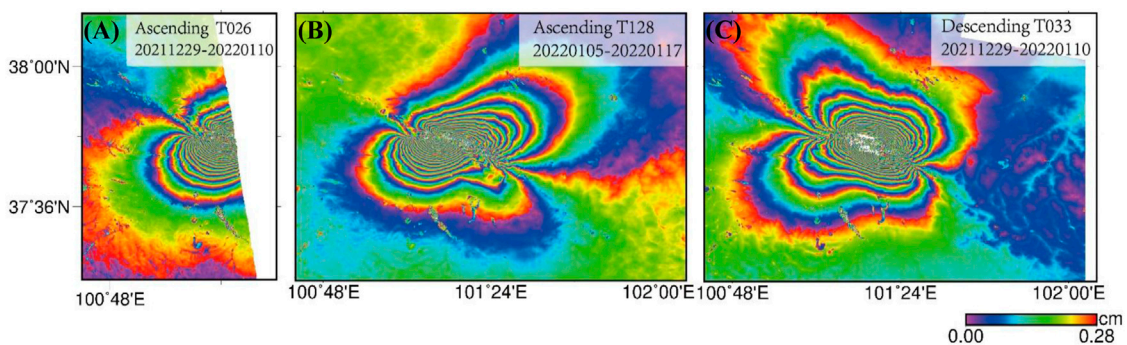


FIGURE 2
Coseismic interferograms of the 2022 Mw6.7 Menyuan earthquake observed by InSAR. (A) Ascending track 026, (B) ascending track 128, and (C) descending track 033 of the Sentinel-1A.

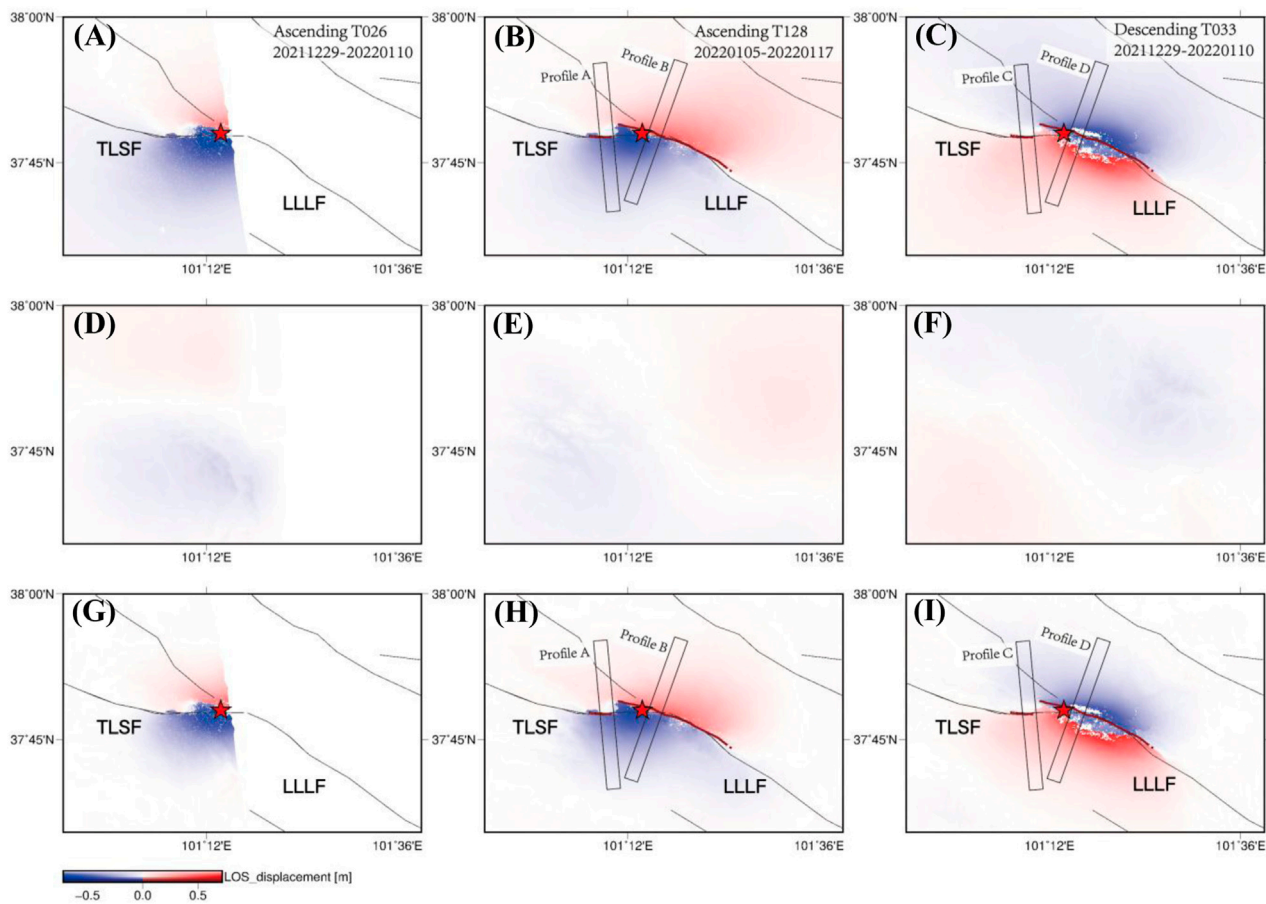
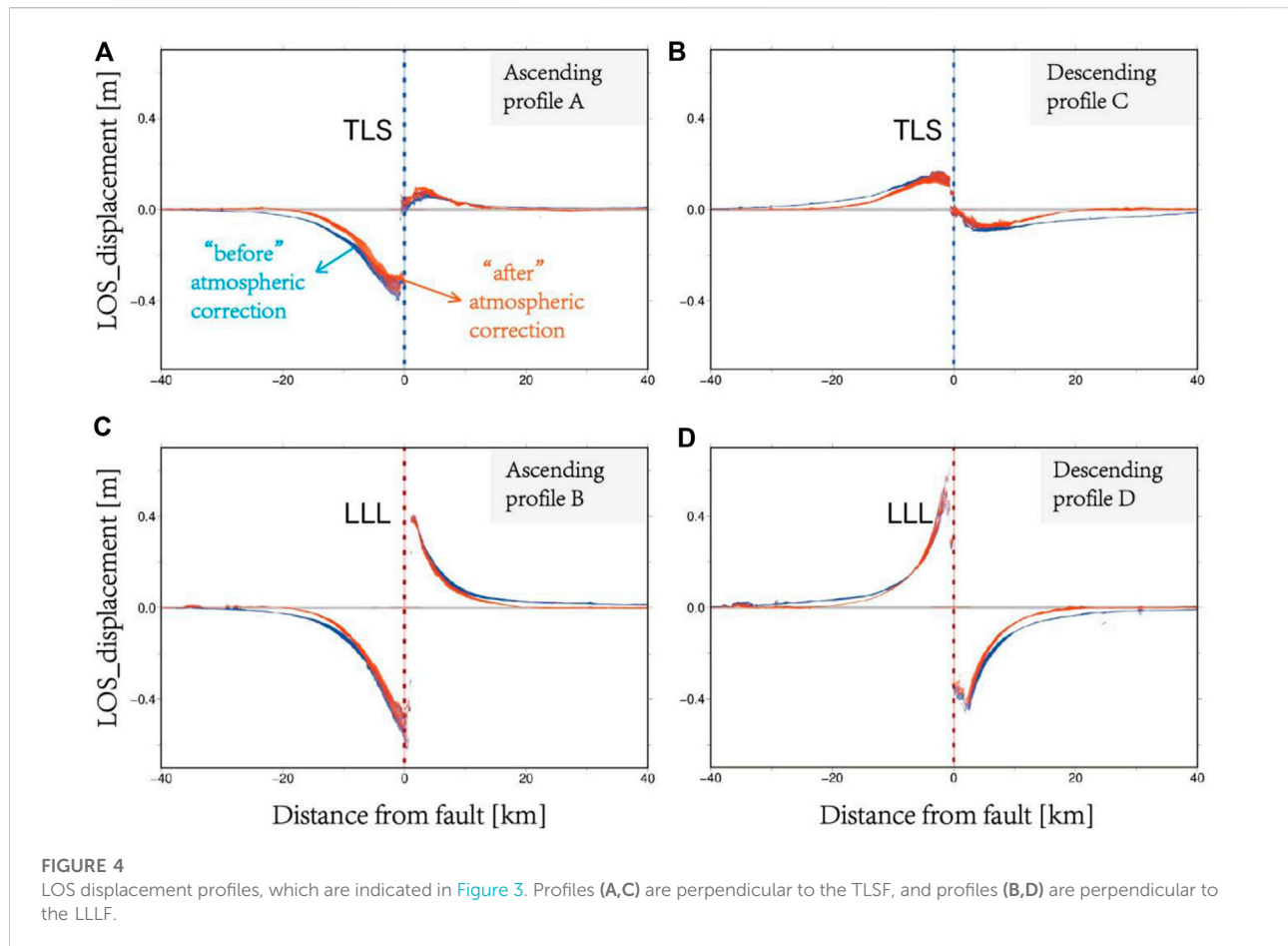


FIGURE 3
Atmospheric correction for the ascending and descending data. (A–C) LOS displacement maps before atmospheric correction. (D–F) Stratified tropospheric delay, estimated using the SSC method. (G–I) After atmospheric correction. The epicenter is indicated by a red star, and the dark-red line indicates the field-investigated surface rupture from Pan et al. (2022).



similar to the estimates in the middle section. In the western section, a slip rate of 4.6 ± 0.2 mm/yr can be seen on the InSAR-inverted results (Huang et al., 2022). No Mw 5+ earthquake was recorded on the LLLF in instrumental record history. Two moderate-magnitude thrust earthquakes occurred on the MDF and northern LLLF in 1986 and 2016, respectively.

Coseismic displacement from InSAR

The C-band TOPS-mode Sentinel-1A images with different observation geometries were used to extract the coseismic displacement maps of the 2022 Menyuan earthquake in the line-of-sight (LOS) direction. The ascending (track 026 and 128) and descending (track 33) interferometric pairs were processed using GAMMA software (Wegmüller et al., 2016). Interferograms were produced with a registration accuracy higher than 0.001 pixel. A 30-m ALOS World 3D DEM was used to remove the topographic phase. The precise orbit data from the European Space Agency (ESA) were employed to reduce the orbital artifacts. After adaptive filtering (Goldstein and Werner, 1998), differential interferograms were unwrapped using the minimum cost flow (MCF) methods

(Werner et al., 2002). The coseismic interferometric fringes in Figure 2 show a deforming area of $\sim 120 \times 80$ km². Similar butterfly-shaped fringe patterns on the descending and ascending interferograms with opposite deformation directions (Figure 2) demonstrate a predominant left-lateral strike-slipping movement on the seismogenic fault of this earthquake. Two cross-fault profiles with 80 km length and 2 km width are extracted from the unwrapping results (Figure 3 and Figure 4). The results reveal that the LOS displacements on the ascending track 128 vary from -0.61 to 0.44 m; they vary from -0.68 to 0.45 m for the ascending track 026 and from -0.51–0.64 m for the descending track 033. The maximum displacement is measured to be ~ 0.68 m. The range offset maps are extracted from the ascending track 128 and descending track 33 Sentinel-1A data (Figure 5). A clear surface rupture can be identified based on them, which fits well with the field investigation of Pan et al. (2022) (Figure 5A).

The uncertainties in InSAR observations have a significant influence on the inversion accuracy of the fault slip models (Lohman and Simons, 2005; Dawson and Tregoning, 2007), especially for small-to-moderate-magnitude earthquakes, whose coseismic surface displacement is small, it is easy to be contaminated by noises in interferograms, such as atmospheric

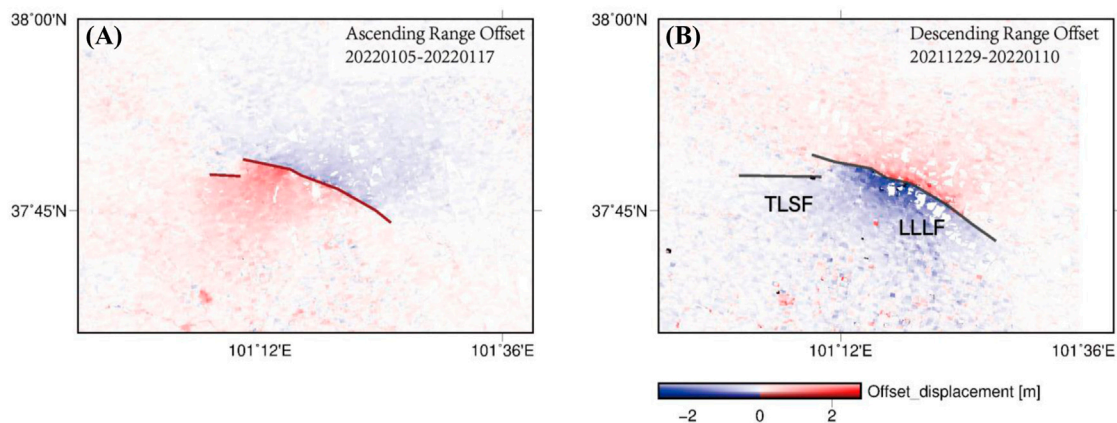


FIGURE 5

Range offsets derived from Sentinel-1A ascending track 128 (A) and descending track 033 (B) by using the offset-tracking technique. The dark-red line marked in (A) indicates the field-investigated surface rupture from Pan et al. (2022), and the gray line marked in (B) indicates the two-segment fault model used in our Slip inversion in Section 4.

noise. Simulation from Dawson and Tregoning (2007) demonstrates coseismic deformation caused by an Mw6.2 earthquake deeper than 10 km, or an Mw 5.5 earthquake deeper than 6 km, is undetectable by the InSAR technique. Therefore, it is necessary to correct the atmospheric signals in interferograms before using them as constraints in slip inversion. Considering a relatively small deforming area generated by small-to-moderate-magnitude earthquakes, the local atmospheric effect on the interferometric phase is mainly from the stratified tropospheric delay. To mitigate the stratified tropospheric effect, a simple-stratification-correction (SSC) approach dedicated to small-to-moderate-magnitude earthquakes was implemented, whose performance has been validated using 23 real earthquakes in Gong et al. (2022). Comparing the ascending and descending results “before” and “after” correction (Figures 3, 4), we can see that the SSC performs well in reducing local stratified contribution in this event. The relative deformation between the hanging wall and footwall is reduced by ~ 2% (~2 cm) (Figure 4).

Slip inversion

To examine the coseismic slip model, the InSAR data from three tracks (26, 128, and 33) and offsets from two tracks (128 and 33) were used to invert for the slip distribution. A homogeneous elastic dislocation model was adopted to calculate the Green’s function using the EDGRN program (Wang et al., 2013), and the Poisson ratio was set to 0.25. We manually cut off the near-fault InSAR data to avoid the influence of unwrapping errors due to the low coherence near the rupture. Three LOS displacement maps and two range offset data are down-sampled using the QuadTree algorithm (Jónsson et al., 2002) to reduce the data volume. Based on the surface displacement gradient on the offset results and field investigation from Pan et al.

(2022), the fault surface trace is mapped, composed of a ~22.7-km-long segment on the western section of LLLF and a ~4.5-km one on the eastern section of the TLSF (Figures 3, 5A). Therefore, we assumed a south dipping two-segment fault model for the inversion (Figure 5B). We fixed the fault trace on the ground for the two segments, and then they are extended down dip and along strike. The fault planes were equally discretized into subfaults with a size of 2 km by 2 km. An efficient FORTRAN program “steepest decent method” (Wang et al., 2013) was applied to resolve the optimal slip model with the weight factors of 1.0 and 0.5 allocated to the InSAR data and offset data, respectively. A grid search method was performed to estimate the fault dips in the range of 50–90° in steps of 2°. Simultaneously, the slip distribution was determined.

To analyze the influence of atmospheric correction on coseismic slip inversion, we conducted inversions twice by using two groups of the InSAR dataset— Group A: the ascending and descending InSAR LOS displacement fields “before” atmospheric correction; Group B: ones “after” atmospheric correction. The InSAR observations, best-fitting models, residuals, and slip distribution from Group B are shown in Figures 6, 7 (see the corresponding results from Group A in Supplementary Figures S2, S3). We found that both models give a good data-model correlation (0.9617 and 0.9535), but the mean slip and moment magnitude is overestimated before atmospheric correction for this event by comparing two inversion results. The mean slips decrease by ~ 12.5% from 0.24 to 0.21 m for the eastern TLSF rupture and by ~ 18.3% from 0.82 to 0.67 m for the western LLLF rupture after correction (Table 3). Accordingly, the moment magnitude decreases from Mw6.74 to Mw6.69 after atmospheric correction, closer to the solutions from GCMT and USGS (Table 1). Both models gave a similar

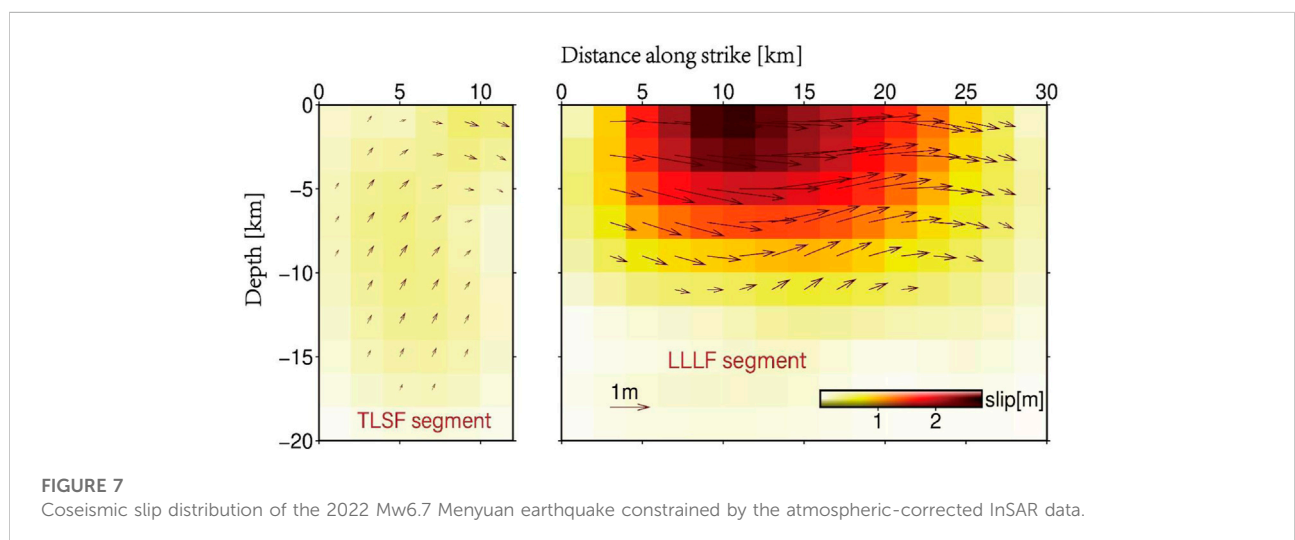
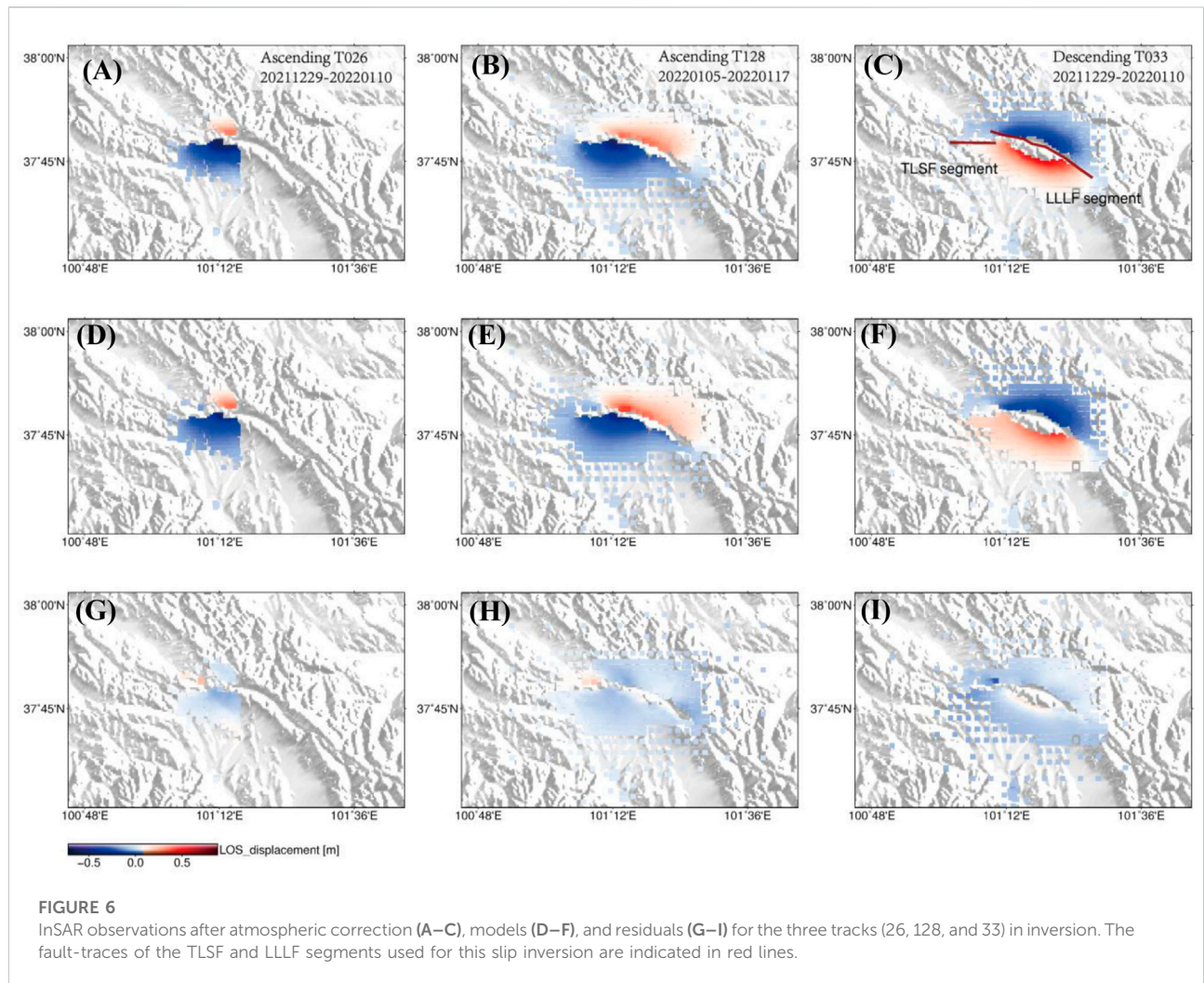


TABLE 3 Inversion results of the 2022 Menyuan earthquake from different data set groups.

Source	Mag. (M_w)	Data-model correlation	Fault segment	Mean strike ($^{\circ}$)	Mean rake ($^{\circ}$)	Mean slip (m)
Group A	6.74	0.9617	TLSF segment	90.89	25.04	0.24
			LLLF segment	114.87	-0.78	0.82
Group B	6.69	0.9535	TLSF segment	90.89	42.16	0.21
			LLLF segment	114.87	0.41	0.67

Group A: the ascending and descending InSAR LOS displacement fields “before” atmospheric correction; Group B: ones “after” atmospheric correction.

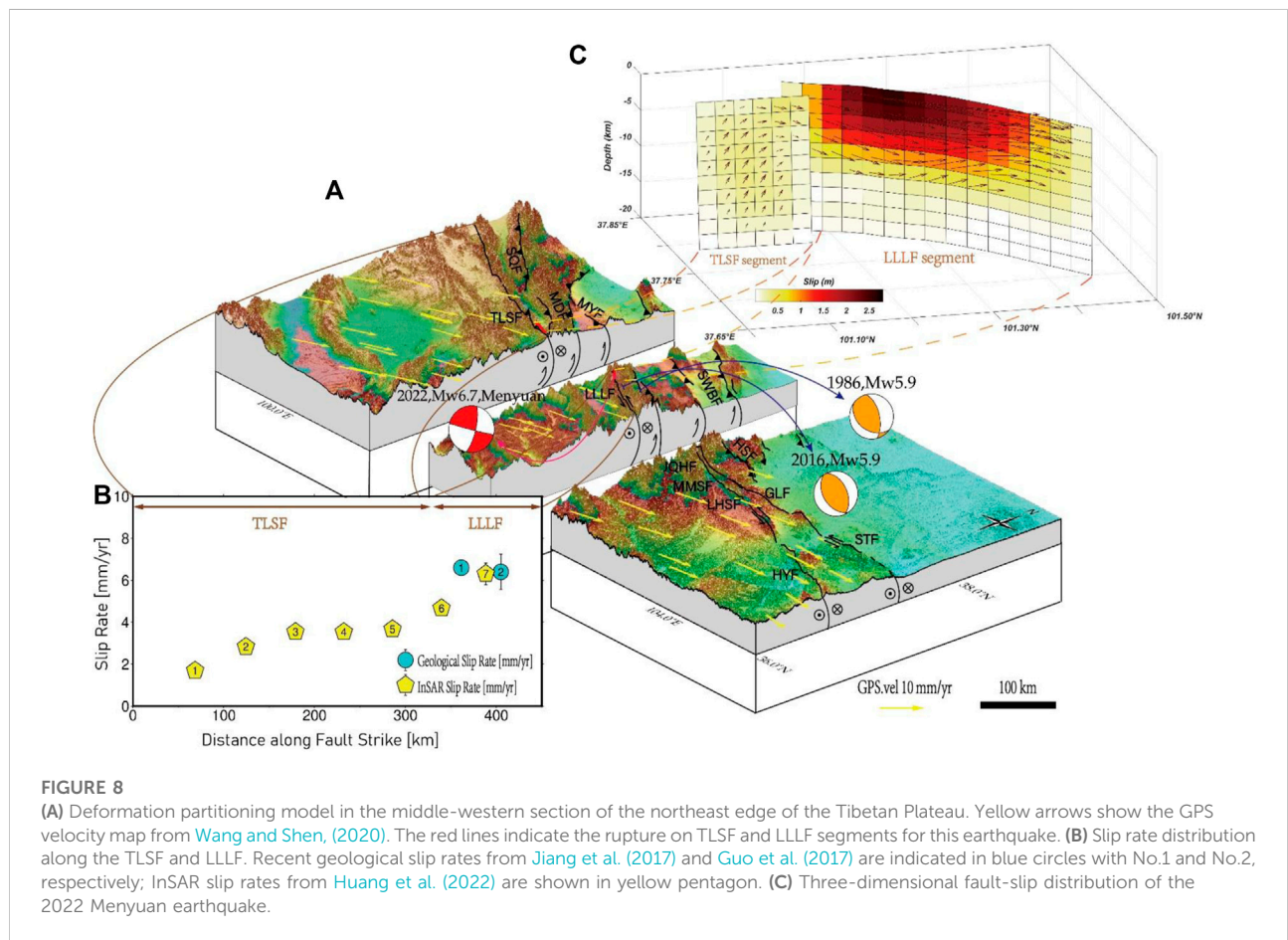


FIGURE 8 (A) Deformation partitioning model in the middle-western section of the northeast edge of the Tibetan Plateau. Yellow arrows show the GPS velocity map from Wang and Shen, (2020). The red lines indicate the rupture on TLSF and LLLF segments for this earthquake. (B) Slip rate distribution along the TLSF and LLLF. Recent geological slip rates from Jiang et al. (2017) and Guo et al. (2017) are indicated in blue circles with No.1 and No.2, respectively; InSAR slip rates from Huang et al. (2022) are shown in yellow pentagon. (C) Three-dimensional fault-slip distribution of the 2022 Menyuan earthquake.

slip distribution pattern with two asperities, located on the western LLLF rupture segment and eastern TLSF rupture segment. The main rupture concentrates on the LLLF segment with the predominant slip patches above 10 km, which takes ~80% of total releasing moment energy. The best-fitting dips for two segments are approximately vertical (88°). In agreement with the focal mechanism solutions, the slip on the LLLF rupture segment is purely left-lateral strike-slip motion with a maximum slip of 2.90 m, which is consistent with the results from Li et al. (2022) and Yu et al. (2022). On the TLSF rupture segment, another asperity with a small slip magnitude is found. In contrast to the left-lateral strike-slip movement on the LLLF segment,

the motion on the TLSF segment is mainly oblique slip at depth, with an obvious thrust component.

Implications for the strain partitioning pattern in the northwestern section of the QL–HYF zone

The modeling results from Bowman et al. (2003) showed that oblique contraction or extension at depth along tectonic boundaries is expected to be accommodated on

predominantly strike-slip and dip-slip faults at the surface, which are observed in many large earthquakes on the block boundaries, such as the 2016 Mw7.0 Kumamoto earthquake in Japan: strike-slip plus normal faults (Toda et al., 2016), the 2018 Mw 7.5 Sulawesi earthquake in Indonesia: strike-slip plus normal faults (Song et al., 2019), and the 1927 M8.0 Gulang earthquake in China: strike-slip plus thrust faults (Guo et al., 2020). As block convergence boundaries in the northeast edge of the Tibetan Plateau, the QL–HYF zone is undergoing oblique compression and strain accumulation, accommodating part of the NE expansion of the Tibetan Plateau together with the GLF, XTF, HSF, and SWBF in the northwestern section of the QL–HYF (west of LLLF–JQHF–GLF triple junction), oblique convergence is partitioned into slip on the purely strike-slipping LLLF and thrusting faults in the north, which is verified by recent findings about the surface rupture of the 1927 Gulang large earthquake: a ~120-km-long strike-slip rupture along the LLLF and a ~42-km-long thrust rupture along the SWBF and eastern HSF (Guo et al., 2020). There are some questions related to such slip partitioning system that remain unknown: 1) where the oblique compression starts to be completely partitioned into left-lateral strike-slip motion and thrust faulting at the ground surface in the northwestern area of the QL–HYF zone and 2) why the same style of deformation partitioning does not maintain along the middle-east section of the QL–HYF (east of LLLF–JQHF–GLF triple junction).

The optimal slip model achieved in Section 4 shows that this earthquake has ruptured two fault segments: the western LLLF segment with pure strike-slip motion and a small part of the eastern TLSF section with an obvious oblique slip (Figure 8C). Two moderate-magnitude thrust earthquakes, the 1986 and 2016 Mw5.9 Menyuan earthquakes, also occurred in the north of LLLF. Combined with the slip-partitioning rupture in the 1927 Gulang earthquake, it is reasonable to suggest that the oblique slip at depths along the whole LLLF segment has completely partitioned into strike-slip and thrust slip at surface in the transpressional QL–HYF systems. As for the TLSF segment, the interseismic InSAR slip rate map shows an obvious step between the eastern TLSF segment (3.5 mm/yr) and LLLF segment (6.4 mm/yr) (Figure 8B). The discrepancy in the left-lateral strike-slip rate between different segments along the QL–HYF is caused by how the strain is partitioned on the QL–HYF and other faults in the north. Given that the oblique slip occurred on the fault plane on the eastern TLSF segment in this earthquake, the strike-slip rate discrepancy between the eastern TLSF segment and LLLF segment is taken up by dip-slip on the TLSF segment together with the shortening of the Qilianshan in the north. Therefore, we suggest that the TLSF–LLLJF junction is a demarcation point where the strain partitioning pattern has changed (Figure 8A). Such transformation is likely to be controlled by the change of fault geometry relative to the direction of regional block movement. Geological investigation shows the TLSF steps left in a ~3-km en

echelon pattern at the western end of LLLF (Guo et al., 2020; Pan et al., 2022). Likewise, in the middle-east section of the QL–HYF zone, the deformation partitioning pattern is determined by the geometric relationship between the fault strike and the region extending direction of the Tibetan Plateau. From the GPS velocity map and fault traces (Figure 1, Figure 8A), the angle between the strike of the JQHF–MMSF–LHSF segment and the direction of block movement is much smaller in comparison to the LLLF segment, especially for the GLF and western section of XTF in the north, the fault strike is almost consistent with the direction of block movement (Figure 1), so the expansion of the Tibetan Plateau is mainly absorbed on these two faults in the form of strike-slip.

Conclusion

This study carried out interferometric processing with a stratified tropospheric correction dedicated to small-to-moderate-magnitude earthquakes for the ascending and descending interferometric pairs from Sentinel-1A satellite to obtain the coseismic deformation maps for the 2022 Menyuan earthquake. The InSAR-inverted results show that the interferometric deformation maps without atmospheric correction will overestimate the mean slip and moment magnitude for this event. The optimal slip model demonstrates that two segments have been ruptured: a >20-km-long section on the western LLLF with purely left-lateral striking-slip motion and a small part of eastern TLSF with an obvious oblique slip, which implies a change in the deformation partitioning pattern at the surface. Recent slip rate studies and the rupture characteristics of historical events also support the same. The change in the strain partitioning pattern along the northeastern edge of the Tibetan Plateau is likely to be controlled by the change of fault geometry relative to the direction of regional block movement (or the regional principal compressive stress direction). Our study is important for improving understanding of the tectonic transformation style in the NE Tibetan Plateau.

Data availability statement

The original contributions presented in the study are included in the article/Supplementary Material; further inquiries can be directed to the corresponding author.

Author contributions

YYT: InSAR data processing, modeling, drawing figures, and data collection. SXG: writing, idea, and data analysis. GWY: atmospheric correction. QCY: suggestion and communication.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2022.1000349/full#supplementary-material>

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