

1 Variations in the light absorption coefficients of  
2 phytoplankton, non-algal particles and dissolved organic  
3 matter in reservoirs across China

4 Yingxin Shang<sup>1,3,5</sup>, Pierre-Andre Jacinthe<sup>2</sup>, Lin Li<sup>2</sup>, Zhidan Wen<sup>1,3</sup>, Ge Liu<sup>1,3</sup>, Lili Lyu<sup>1,5</sup>, Chong  
5 Fang<sup>1,5</sup>, Bai Zhang<sup>1,3</sup>, Junbin Hou<sup>1</sup>, Kaishan Song<sup>1,3,4\*</sup>

6 <sup>1</sup> Northeast Institute of Geography and Agroecology, CAS, Changchun, 130102, Jilin, China.

7 <sup>2</sup> Department of Earth Sciences, Indiana University–Purdue University, Indianapolis, IN, USA.

8 <sup>3</sup> Changchun Jingyuetan Remote Sensing Observation Station, Chinese Academy of Sciences,  
9 Changchun 130102, China

10 <sup>4</sup> School of Environment and Planning, Liaocheng University, Liaocheng, 252000, China

11 <sup>5</sup> University of Chinese Academy of Science, Beijing, 100049, China.

12 Corresponding author: Kaishan Song ([songkaishan@iga.ac.cn](mailto:songkaishan@iga.ac.cn))

13 Key Points:

- 14 ● Absorption parameters varied significantly from different regions and trophic states
- 15 ● Factors affecting the three inherent optical absorption for large-scale regions were explored
- 16 ● Dominant absorption types for various regions were determined

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18 **Abstract**

19           Reservoirs were critical sources of drinking water for many large cities around the world,  
20 but progress in the development of large-scale monitoring protocols to obtain timely information  
21 about water quality had been hampered by the complex nature of inland waters and the various  
22 optical conditions exhibited by these aquatic ecosystems. In this study, we systematically  
23 investigated the absorption coefficient of different optically-active constituents (OACs) in 120  
24 reservoirs of different trophic states across five eco-regions in China. The relationships were  
25 found between phytoplankton absorption coefficient at 675 nm ( $a_{ph}(675)$ ) and Chlorophyll  
26 *a* (*Chla*) concentration in different regions ( $R^2$ :0.60-0.82). The non-algal particle (NAP)  
27 absorption coefficient ( $a_{NAP}$ ) showed an increasing trend for reservoirs with trophic states.  
28 Significant correlation ( $p<0.05$ ) was observed between chromophoric dissolved organic matter  
29 (CDOM) absorption and water chemical parameters. The influencing factors for contributing the  
30 relative proportion of OACs absorption including the hydrological factors and water quality  
31 factors were analyzed. The non-water absorption budget from our data showed the variations of  
32 the dominant absorption types which underscored the need to develop and parameterize region-  
33 specific bio-optical models for large-scale assessment in water reservoirs.

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## 38 **1 Introduction**

39           The OACs (optically-active constituents) in inland waters typically included  
40 phytoplankton, NAP and CDOM (Ylöstalo et al., 2014). The inherent optical properties (IOPs)  
41 of OACs in water bodies were depend on the medium and independent of the ambient light field,  
42 so that underwater light absorption coefficients of OACs can be used to investigate the  
43 variability of OACs in natural waters under various aquatic and environmental conditions (Lyu  
44 et al., 2020, Matsuoka et al., 2011). The complexity of sources and composition of OACs  
45 enhance the variability of optical properties of Case-2 waters, including inland water reservoirs  
46 (Li et al., 2017, Wu et al., 2011). Reservoirs were constructed to meet a variety of needs,  
47 including irrigation, recreation and drinking water storage, and consequently contribute to the  
48 survival and well-being of human communities. Water quality in reservoirs can be a challenge,  
49 thus creating the need to establish monitoring programs that are flexible, reliable and cost-  
50 effective, and can deliver timely information about water quality status (Shi et al., 2019a).  
51 According to our comparisons of optical absorption coefficients of reservoirs and lakes across  
52 five eco-regions based on in situ field sampling dataset, the mean optical coefficients at 440nm  
53 including CDOM, phytoplankton and non-algal particulates for reservoirs are significant higher  
54 that of reservoirs ( $p < 0.05$ ) (Table S1). There was great interest in understanding light absorption  
55 process of OACs of reservoirs in order to develop water quality monitoring programs that  
56 provide accurate information about *in-situ* conditions and ensure a safe water supply (Shang et  
57 al., 2021).

58           The *in situ* study for the linkage of optical absorption and water quality parameters was  
59 useful to identify the quality and quantity of the components and further calibrated and validated  
60 remote sensing models of water quality parameters (Gholizadeh et al., 2016; Zhao et al., 2017).

61 Phytoplankton absorption coefficients have been related to phytoplankton biomass and *Chla*  
62 (Binding et al., 2008). NAP generally consisted of sediments, non-algal organic detritus and  
63 living non-algal particulates (Binding et al., 2008). The CDOM also affected the optical  
64 properties of aquatic systems through control of UV and blue radiation penetration the water  
65 column (Organelli et al., 2016). CDOM was generally derived from autochthonous and  
66 allochthonous sources (Jones et al., 2009). Generally, CDOM absorption coefficients at 355nm  
67 and 440nm was chosen to represent the relative absorption concentration, and used as CDOM  
68 absorption parameters to build remote sensing retrieval for inland waters (Brezonik et al., 2015).  
69 Except for absorption coefficient, the absorption slope can provide information about the source  
70 and composition of CDOM (Helms et al., 2008; Zhang et al., 2011).

71         The specific absorption coefficients of OACs could reflect regional variability in inherent  
72 optical characteristics of aquatic systems as well as their ecological conditions and trophic state  
73 (Shi et al., 2019b; Shang et al., 2019). The specific phytoplankton absorption coefficients can  
74 significantly vary among the IOPs of various substances in aquatic systems (Wan et al., 2020;  
75 Bricaud et al., 2010; Sun et al., 2012). The intensity and variability of light absorption can be  
76 related to the trophic state of a water body since the average size of phytoplankton cells usually  
77 increases from oligotrophic to eutrophic waters (Shi et al., 2019b; Ciotti et al., 2002).  
78 Investigations into the specific absorption coefficient of NAP have also shown that this  
79 parameter varies with region and season (Sun et al., 2012; Tilstone et al., 2012; Astoreca et al.,  
80 2012). Particle sizes and composition were key factors determining the optical absorption and  
81 scattering characteristics of NAPs in areas such as coastal waters, continental shelves, and semi-  
82 arid regions (Binding et al., 2005; Wen et al., 2016). Previous studies have shown that absorption  
83 slopes can vary widely depending on the eco-region of the water body investigated, with high

84 slopes of CDOM generally observed in water bodies from high-elevation regions likely due to  
85 strong photo-chemical processes (Brito et al., 2015; Babin et al., 2003; Zhang et al., 2011). Due  
86 to the complex composition of OACs encountered in reservoirs of inland waters, and the variable  
87 hydrological (reservoir's storage) and geomorphological environments present in these  
88 waterbodies, light absorption characteristics vary significantly (Li et al., 2012; Vantrepotte et al.,  
89 2012).

90         In many countries around the world, including China, reservoirs account for a significant  
91 part of inland waters. As reservoirs are often primary sources for agricultural, urban and  
92 industrial uses, they are critical for economic and social development. Previous studies  
93 addressing these limitations have generally focused on one reservoirs; consequently, results of  
94 these past studies lack regional perspectives and were inadequate to provide a wider view on the  
95 variability of OACs across eco-regions (Li et al., 2012; Naik et al., 2013; Shi et al., 2013).  
96 According to the Ecological Environment Bulletin in 2017 from Ministry of Ecology and  
97 Environment of the People's Republic of China, 23% of the monitoring lakes or reservoirs were  
98 eutrophic and 67% of the monitoring important reservoirs or lakes were mesotrophic. Therefore,  
99 a comprehensive study of light absorption characteristics in reservoirs spanning across several  
100 eco-regions and trophic states would be needed to provide the foundation for water quality  
101 monitoring using optical models. The objectives of this study were to: (1) examine the spatial  
102 variation of OAC absorption with respect to eco-region and trophic state of water reservoirs; (2)  
103 analyze the factors affecting the three inherent optical absorption for large-scale regions were  
104 explored; and (3) the dominant absorption types for various regions were determined

## 105 **2 Materials and Methods**

### 106 **2.1 Study Area**

107           Reservoirs as important drinking water sources according to the list from the Ministry of  
108 Ecology and Environment of China were selected to represent the diversity of geographical  
109 locations, morphological characteristics, ecological conditions and water uses. For this study, we  
110 selected 120 reservoirs distributed across five regions in China. Regions were defined on the  
111 basis of variation in hydrological and geological conditions (Zhao et al., 2010), and included: the  
112 Northeast hydrological region (NLR), the Eastern hydrological region (ELR), the Yungui plateau  
113 hydrological region (YGR), the Mengxin hydrological region (MXR) and the Tibetan Plateau  
114 hydrological region (TPR) (Fig. 1). Reservoirs ( $>10^4 \text{ m}^3$ ) as important drinking water sources  
115 across China were selected to represent the diversity of geographical locations, morphological  
116 characteristics, ecological conditions and water uses. Accessibility was an additional factor  
117 considered in the selection of water reservoirs for this study. All the reservoirs selected were  
118 relatively large and representative of the five regions described above.

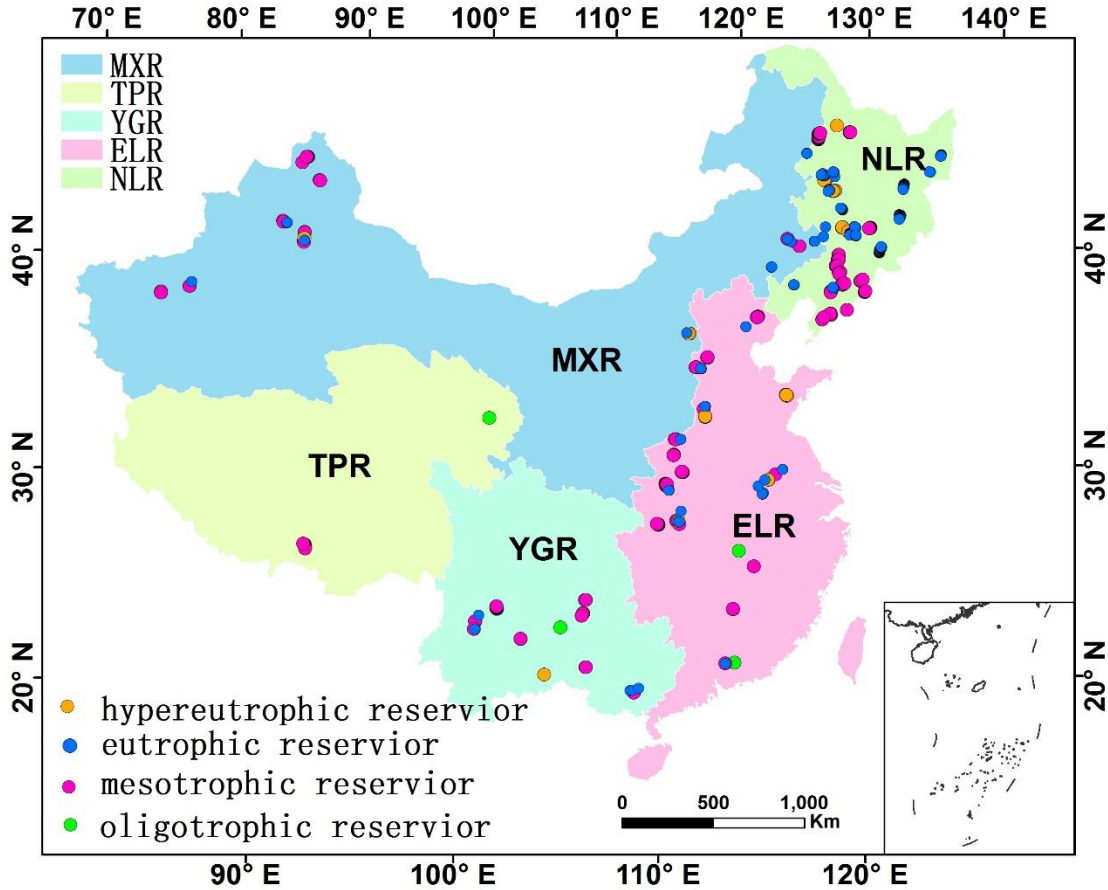


Figure 1 The distribution of sampling reservoirs and main regions across China.

## 2.2 Sampling description

A total of 507 water samples were collected from 120 reservoirs (3-5 samples in each reservoir; Fig.1) across the five regions of China in autumn between 2014 and 2015 with the consideration of diversity of geographical locations, morphological characteristics, ecological conditions and water uses. Sampled reservoirs exhibit a range of trophic states. Due to the uneven distribution of numbers of reservoirs among regions and the convenience of taking fieldtrips, more water samples were collected in the NLR (40 reservoirs) and ELR regions (39 reservoirs) than in the TPR (3 reservoirs), YGR (17 reservoirs) and MXR regions (21 reservoirs). Water samples were generally collected from the central part of the reservoir. Surface water

131 samples (0.5 m) were collected in 1 L acid-washed plastic bottles, transported in a portable  
132 refrigerator, and filtered within 24 h upon returning to the laboratory. Water quality parameters  
133 such as temperature, dissolved oxygen, salinity, pH, turbidity, ORP, *Chla*, fluorescent colored  
134 dissolved organic matter were measured using in-situ multi-parameter probes (YSI 600 XLM  
135 Sonde, YSI Inc., Yellow Springs, OH). Secchi disk depth (SDD) was measured with a Secchi  
136 disk.

### 137 **2.3 Laboratory Measurements**

138 Water samples were initially filtered through a 0.45  $\mu\text{m}$  cellulose acetate micro-porous  
139 membrane filter (Peninsula). Dissolved organic carbon (DOC) concentration of water samples  
140 was determined using a Shimadzu TOC-5000 analyzer. The concentration of *Chla* was  
141 determined by analyzing the absorption spectra of samples extracted with 90% acetone for 24 h  
142 at 4 °C using a Shimadzu UV-2600PC spectrophotometer (Shang et al., 2019). To determine the  
143 concentration of total solid matter (TSM), water samples were filtered through pre-weighed 47 mm,  
144 0.7  $\mu\text{m}$  pore size glass fiber filters which were then oven dried at 85 °C for at least two hours  
145 before cooling and re-weighing. The filters were again heated in a muffle furnace at 400 °C for  
146 two hours (Wen et al., 2016) to determine the concentration of MSS (mineral suspended solids)  
147 by re-weighing the filters after ignition. Total phosphorus (TP) were measured after oxidation in  
148 the presence of boric acid and sodium hydroxide with a standard procedure  
149 (APHA/AWWA/WEF, 1998).

150

### 151 **2.4 OACs absorption measurements**

152 CDOM was extracted by filtering water samples through a 0.22  $\mu\text{m}$  polycarbonate  
153 membrane (Whatman, 110606). The detailed measurement of CDOM spectra and the coefficient



154 calculation process is shown in the supplementary materials. The spectral slopes ( $S_{\text{CDOM}280-400}$   
155 and  $S_{\text{CDOM}275-295}$ ) were determined from least- squares regression of log-transformed  
156 absorption coefficients in the ranges 280-400 nm and 275-295 nm (Binding et al., 2008; Helms  
157 et al. 2008; Song et al., 2010). The specific CDOM absorption at 355nm ( $a^*_{\text{CDOM}}$ ) which was  
158 calculated by  $a_{\text{CDOM}}(355)$  normalized to DOC concentration was obtained as well.

159 Total particulate absorption coefficients ( $a_p$ ) were determined using the quantitative filter  
160 technique (QFT) (Mitchel et al., 2002). Water samples were filtered through 0.7  $\mu\text{m}$  pore size  
161 Whatman GF/FTM filters before measuring the absorption coefficients of total particulate  
162 absorption ( $a_p(\lambda)$ ). Non-algal particulate absorption ( $a_{\text{NAP}}(\lambda)$ ) was obtained using the methanol  
163 extraction method to remove pigment matter. The spectral absorption coefficient for NAP was  
164 computed (Babin et al., 2003). The phytoplankton absorption coefficient was then obtained as  
165 the difference between  $a_p$  and  $a_{\text{NAP}}$ . The spectral slope of the  $a_{\text{NAP}}$  ( $S_{\text{NAP}}$ ) was related to particle  
166 size and the relative proportion of mineral and organic particles (Binding et al., 2008). The  $a^*_{\text{NAP}}$   
167 was calculated with the ratio of  $a_{\text{NAP}}$  and MSS (Babin & Stramski, 2004). The specific  
168 absorption ( $a^*_{\text{ph}}$ ) was calculated with the ratio of  $a_{\text{ph}}$  and *Chla* concentrations (We use 440nm  
169 and 675nm) (Bricaud et al., 1995). The detailed calculation of the particulate absorption is shown  
170 in supplementary materials.

171 Total non-water absorption referred to the total absorption of CDOM, phytoplankton and  
172 non-algal particulates (Wen et al., 2016). The relative contribution ratio of the three sources of  
173 absorption was calculated with the light absorption coefficient at 440nm normalized to the total  
174 non-water absorption at 440nm in this study (Li et al., 2015; Wen et al., 2016). The detailed  
175 measurement steps were shown in the supplementary materials.

176

## 177 **2.5 Trophic states assessment**

178 Trophic state assessment of the reservoirs was based on the modified Carlson's trophic  
179 state index ( $TSI_M$ ) (Zhang et al., 2018; Lyu et al., 2020, Aizaki, 1981). This index was  
180 calculated using three limnological parameters, chlorophyll a ( $Chla$ ,  $\mu\text{g/L}$ ), Secchi disk  
181 transparency (SDD, m) and total phosphorus (TP,  $\mu\text{g/L}$ ), according to the following equations:

$$182 \quad TSI_M(Chla) = 10(2.46 + \frac{\ln Chla}{\ln 2.5}) \quad (4)$$

$$183 \quad TSI_M(SDD) = 10(2.46 + \frac{3.69 - 1.53 \ln SDD}{\ln 2.5}) \quad (5)$$

$$184 \quad TSI_M(TP) = 10(2.46 + \frac{6.71 + 1.15 \ln TP}{\ln 2.5}) \quad (6)$$

185 Comprehensive TSI ( $TSI_M$ ) was calculated with the mean index of the three parameters, 0.297  
186 for SDD, 0.54 for  $Chla$  and 0.163 for TP. Results from this calculation range from 0 to 100 and  
187 provide a scale to rate the trophic state of the reservoirs:  $0 < TSI_M \leq 30$  oligotrophic;  $30 < TSI_M \leq 50$   
188 mesotrophic;  $TSI_M > 50$  eutrophic (Shang et al., 2019), among the eutrophic states, the index  
189 larger than 70 was considered as hypereutrophic states.

## 190 **2.6 Hydrological dataset collection**

191 The hydrological dataset including design flood level (the highest water level in a reservoir  
192 after flooding), total storage, flood regulation storage capacity, normal pool level, spillway  
193 design flood level, beneficial reservoir capacity, flood limit level, inactive storage capacity (the  
194 lowest water level under normal operation of a reservoir), dead water level for main reservoirs  
195 was purchased from the local hydrographic offices and the data was used to analyze the  
196 influences of hydrological factors on OACs properties.

## 197 **2.7 Data analysis**

198 The variability of OACs properties in water reservoirs from different regions and trophic  
199 states was examined, and statistical analyses (mean value, standard deviation, linear or non-  
200 linear regression, t-test) were conducted using SPSS 22.0 (SPSS, Chicago, IL, USA). K-means  
201 clustering analysis was used to evaluate the water classification types based on light absorption  
202 coefficients (Spyrakos et al., 2017). Correlation plots of different OACs-related parameters  
203 were constructed using Origin 9.0 (Origin Lab, Northampton, MA). Spatial distribution of  
204 sampling sites was conducted using ArcGIS 10.1 software. Difference among water reservoirs  
205 in regard to OACs indices was assessed with one-way ANOVA. Statistically significant  
206 difference was determined at  $p < 0.01$ .

## 207 **3 Results**

### 208 **3.1 Biogeochemical parameters**

209 For some of the water quality parameters used for trophic state assessment, mean values can  
210 increase by an order of magnitude between oligotrophic and hyper eutrophic reservoirs (Table 1).  
211 The temperature ranged from 10.8 to 22.5°C. No significant difference was observed among  
212 regions ( $p > 0.05$ ). There were significant differences between reservoirs in the MXR and YGR  
213 regions in regard to DOC concentration ( $p < 0.05$ ). TSM concentration ranged from 0.17 to  
214 478.50 mg/L for reservoirs across regions (Table 2). *Chla* concentration ranged from 0.04 to  
215 3380.45  $\mu\text{g}/\text{m}^3$  for different trophic states. *Chla* concentration in different hydrological regions  
216 was different ( $p < 0.05$ ). *Chla* in reservoirs from the ELR, MXR and NLR regions was  
217 significantly higher than in those from the TPR and YGR regions. Overall, water chemistry  
218 parameters in NLR, ELR and MXR water reservoirs were relatively higher than in those from the  
219 TPR and YGR regions.

220

Table 1. Water chemical parameters of different trophic reservoirs

	Oligotrophic (n=4)			Mesotrophic (n=50)			Eutrophic (n=54)			Hyper Eutrophic (n=12)		
	Mean	Min.	Max.	Mean	Min.	Max.	Mean	Min.	Max.	Mean	Min.	Max.
TSM(mg/L)	0.46	0.01	1.80	3.93	0.17	478.50	25.24	1.50	184.00	48.6	14.40	310.00
SDD (m)	5.74	5.08	11.30	2.03	0.36	3.28	0.59	0.01	2.99	0.24	0.10	0.48
<i>Chla</i> (µg/L)	0.33	0.04	1.40	3.81	0.00	107.51	21.65	0.00	61.83	106.03	2.80	3380.45
DOC(mg/L)	2.04	0.67	2.94	2.93	0.70	22.62	8.49	0.91	60.60	10.9	4.03	71.04
TP (mg/L)	0.001	0.0008	0.02	0.03	0.002	0.11	0.06	0.01	1.65	0.58	0.01	10.37

221 Note: TSM=concentration of total solid matter, DOC=dissolved organic carbon, SDD=Secchi disk depth,  
 222 *Chla*=chlorophyll-a concentration; TP=Total Phosphorus; N=numbers of reservoirs

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Table 2 The mean water chemistry parameters for reservoirs in five lake regions.

	NLR		ELR		MXR		TPR		YGR	
	Mean	Min-Max	Mean	Min-Max	Mean	Min-Max	Mean	Min-Max	Mean	Min-Max
TSM(mg/L)	25.79	0.17-478.50	9.92	0.2-310.0	25.28	2.33-134.54	1.01	0.01-4.50	3.53	0.20-19.00
SDD(m)	0.93	0.01-4.32	4.94	0.16-6.80	0.77	0.25-1.70	6.71	3.08-11.30	1.98	0.40-5.80
<i>Chla</i> (µg/L)	22.40	1.02-219.70	46.71	0.31-3380.5	15.63	0.17-213.88	0.41	0.041-1.40	6.91	0.38-22.07
DOC (mg/L)	6.44	0.98-33.56	3.97	0.70-71.04	9.87	1.12-60.60	5.22	1.94-9.46	1.90	0.76-6.29
TP(mg/L)	0.07	0.006-0.833	0.22	0.001-10.37	0.04	0.008-0.108	0.04	0.008-0.114	0.004	0.001-0.015

226 Note: TSM=concentration of total solid matter, DOC=dissolved organic carbon, SDD=Secchi disk  
 227 depth, *Chl-a*=chlorophyll-a concentration; TP=Total Phosphorus;

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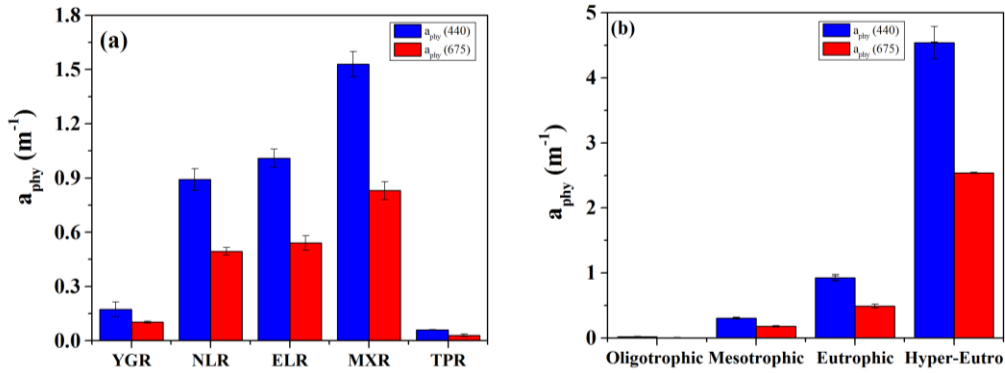
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### 233 3.2 Phytoplankton absorption

234 Across five regions, the average  $a_{ph}(440)$  ranged from  $0.174 \text{ m}^{-1}$  to  $1.53 \text{ m}^{-1}$ , and the  
 235 average  $a_{ph}(675)$  ranged from  $0.103 \text{ m}^{-1}$  to  $0.83 \text{ m}^{-1}$  (Fig. 2a). With respect to  $a_{ph}(440)$ , there were  
 236 significant differences ( $p < 0.01$ ) between the YGR (NLR) and MXR regions, and also between  
 237 trophic states ( $p < 0.01$ ). From oligotrophic to hyper eutrophic reservoirs, the average value for  
 238  $a_{ph}(440)$  ranged from  $0.31 \text{ m}^{-1}$  to  $4.54 \text{ m}^{-1}$ , and  $a_{ph}(675)$  ranged from  $0.18 \text{ m}^{-1}$  to  $2.54 \text{ m}^{-1}$  (Fig.  
 239 2b). In the ELR, YGR and NLR regions, strong relationships ( $R^2$ : 0.73-0.82) were found between  
 240 phytoplankton absorption coefficient at 675 nm and *Chla* concentration (Fig.3). However,

241 relationships between these variables were weaker ( $R^2$ : 0.60-0.68) in the TPR and MXR regions.

242 The value of  $a_{ph}^*$  (440) for various regions and trophic states was shown in Fig.S1.

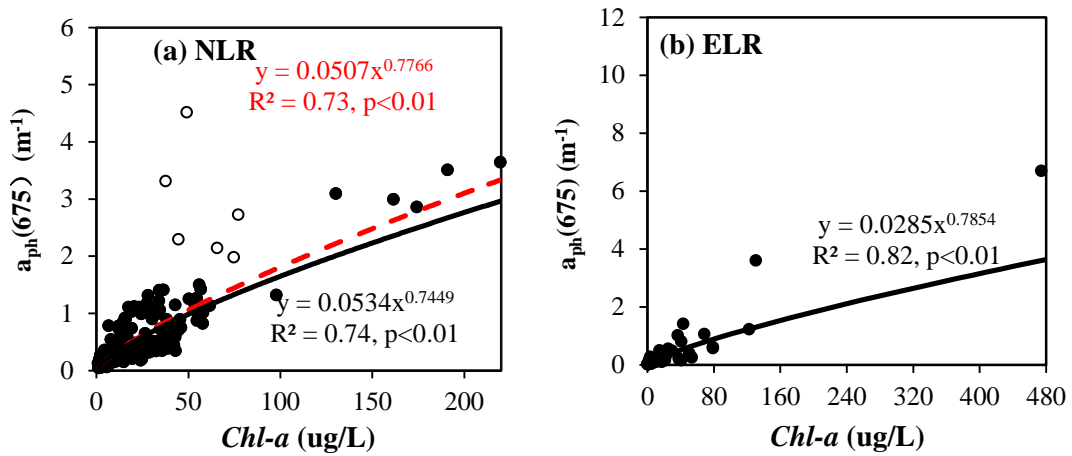


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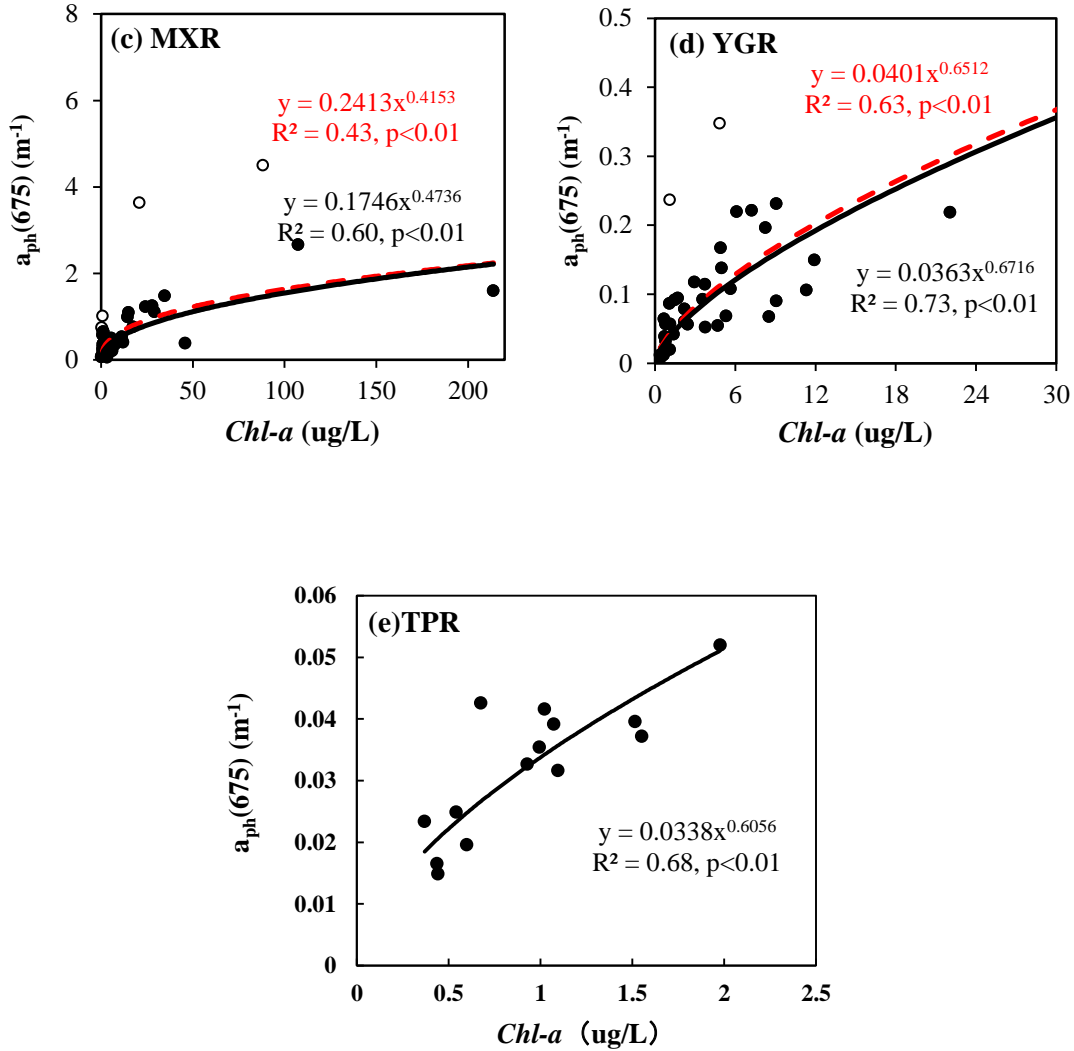
244 Fig.2 (a) Phytoplankton absorption at 440nm and 675 nm for different lake regions and (b)

245 Comparison for  $a_{ph}$  for different trophic status.

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250 Fig. 3 The  $a_{ph}(675)$  plotted against  $Chl-a$  concentration. (a) NLR; (b) ELR; (c)MXR; (d)YGR;  
 251 (e)TPR.

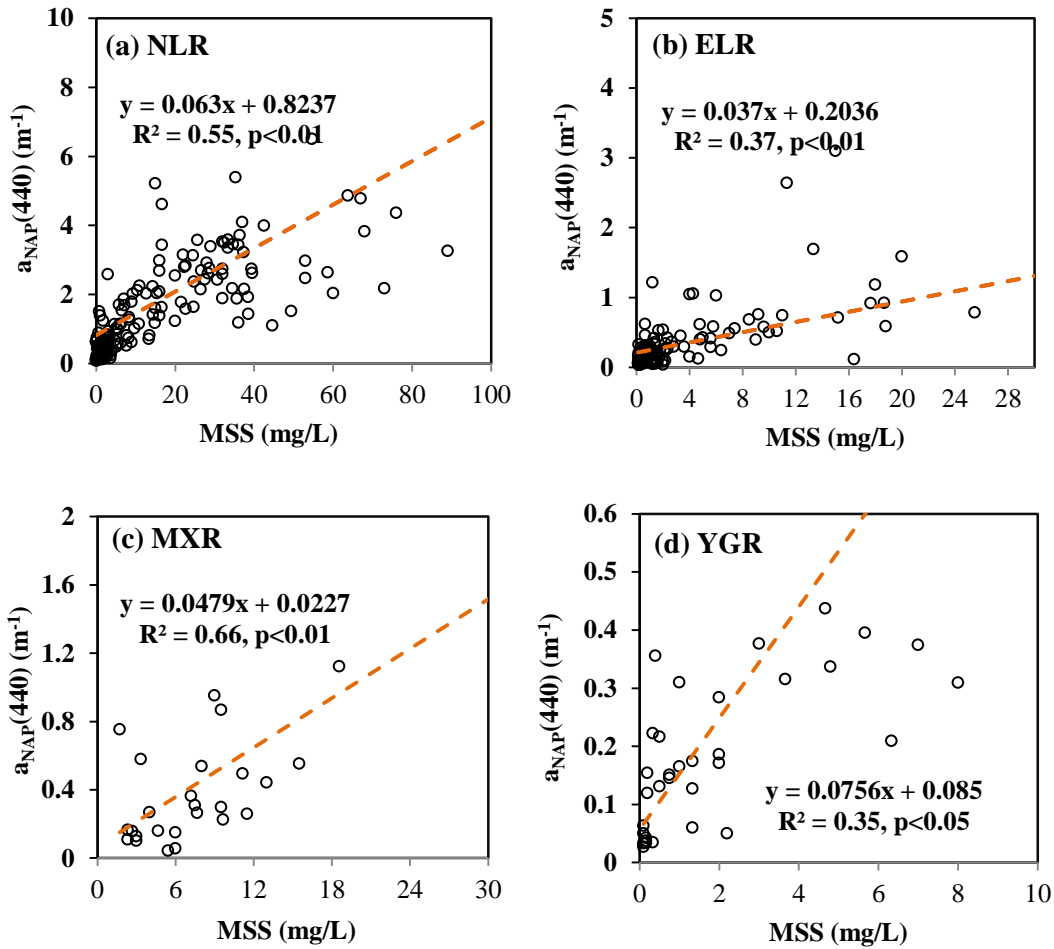
252 Note: The hollow points represented the outliers in different regions; The dotted lines and black lines  
 253 were the exponential models with outliers and without outliers respectively; The correlation equation with  
 254 red was obtained including outliers; The correlation equation with black was obtained for eliminating the  
 255 outliers.

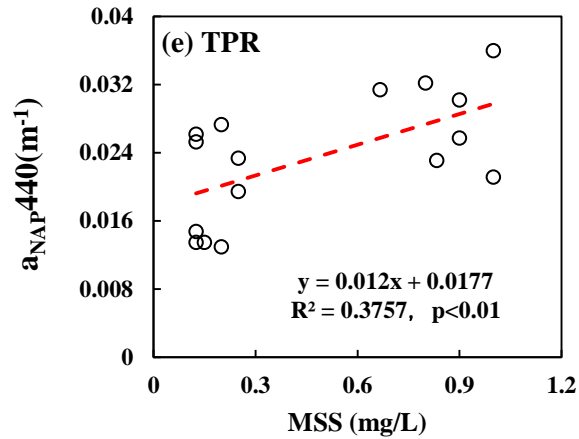
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### 257 3.3 NAP Absorption

258 For all the reservoirs investigated, the absorption coefficients for non-algal particle  
 259 ( $a_{NAP}(440)$ ) had a mean value of  $1.42 \text{ m}^{-1}$ . The non-algal particle absorption coefficient values  
 260 showed an increasing trend for reservoirs with oligotrophic, mesotrophic to hyper-eutrophic

261 states (Fig. S2). The mean  $a_{\text{NAP}}^*(440)$  ranged from 0.021 to 0.43  $\text{m}^2/\text{g}$  for various regions and  
262 trophic states was shown in Fig S3. There were significant differences between YGR and MXR  
263 (TPR) for  $a_{\text{NAP}}^*(440)$ . Significant but not so strong relations ( $p < 0.01$ ,  $R^2$ : 0.35-0.66) were found  
264 between  $a_{\text{NAP}}$  and MSS in different regions (Fig.4). The mean  $S_{\text{NAP}}$  was 0.009, 0.015, 0.011,  
265 0.016 and 0.021  $\text{um}^{-1}$  for NLR, ELR, MXR, TPR and YGR respectively.





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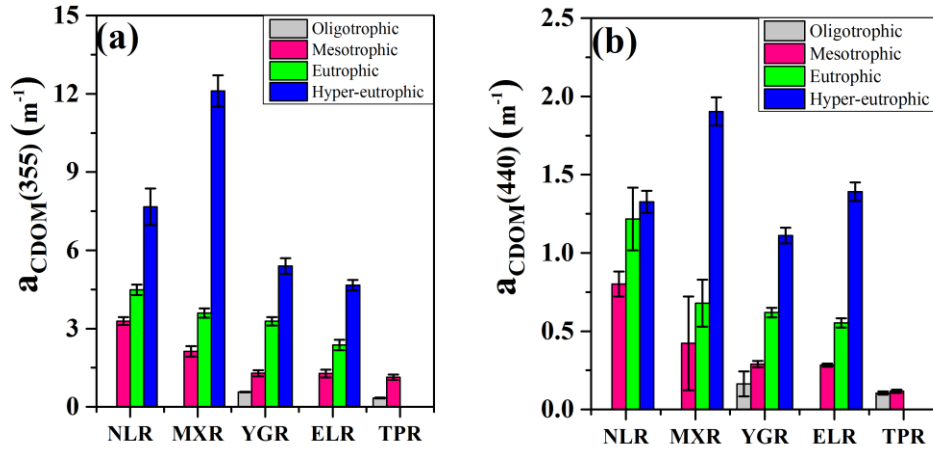
269 Fig. 4 Correlation between NAP absorption and MSS for lake regions, (a) NLR; (b) ELR;  
 270 (c)MXR; (d)YGR; (e)TPR.

### 271 3.4 CDOM absorption

272 The CDOM absorption at 355 nm ranged from 0.34 to 12.11  $\text{m}^{-1}$ , and the CDOM  
 273 absorption at 440 nm ranged from 0.15 to 1.90  $\text{m}^{-1}$  (Fig. 5). There was significant difference  
 274 ( $p < 0.01$ ) among reservoirs of different trophic states within the same region with respect to  
 275 CDOM absorption at 355nm. Within a region, the mean value of  $a_{\text{CDOM}(355)}$  and  $a_{\text{CDOM}(440)}$  for  
 276 the reservoirs significantly increased with trophic state ( $p < 0.05$ ). The  $S_{\text{CDOM}280-400}$  ranged from  
 277 10.99 to 25.47  $\mu\text{m}^{-1}$ , with a mean value of 19.26  $\mu\text{m}^{-1}$  (Fig.6). A decreasing trend for  $S_{280-400}$  and  
 278  $S_{275-295}$  was also observed with the increasing of trophic states (Fig.6). The mean  $a^*_{\text{CDOM}}$  for all  
 279 reservoirs ranged from 0.14 to 0.92  $\text{L mg/m}$  (Fig. S4). With the increasing of trophic states, the  
 280  $a^*_{\text{CDOM}}$  increased significantly ( $p < 0.05$ ). The relationships between  $a^*_{\text{CDOM}(355)}$  and  $S_{280-400}$  was  
 281 significant ( $p < 0.01$ ) (Fig. S5) and the correlations between CDOM absorption coefficients  
 282 ( $a_{\text{CDOM}(440)}$  and  $a_{\text{CDOM}(355)}$ ) and  $S_{275-295}$  for various trophic states and regions were significant  
 283 as well ( $p < 0.01$ ) (Table S2).



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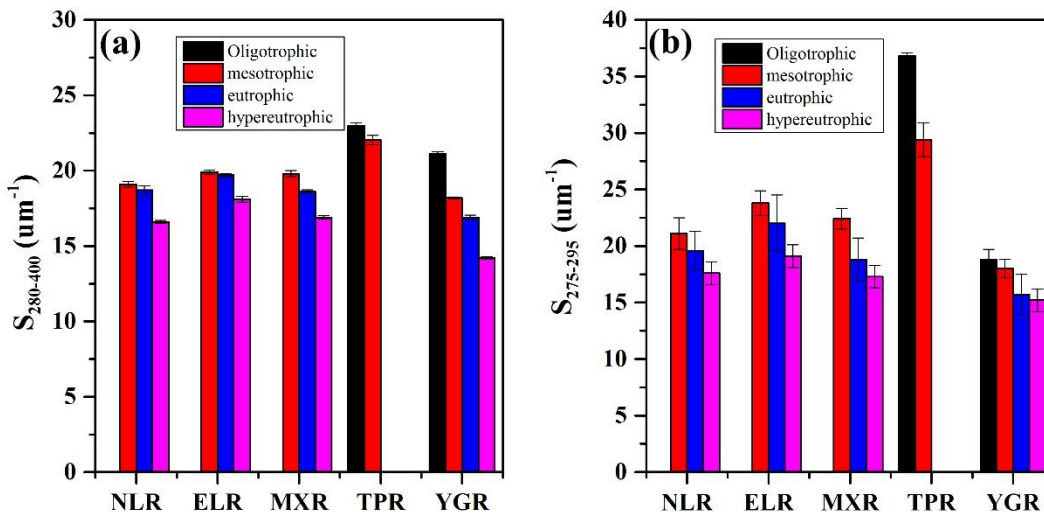
286 Fig.5 The comparisons of CDOM absorption at (a) 355nm, and (b) 440nm in different regions

287 concerning with various trophic states.

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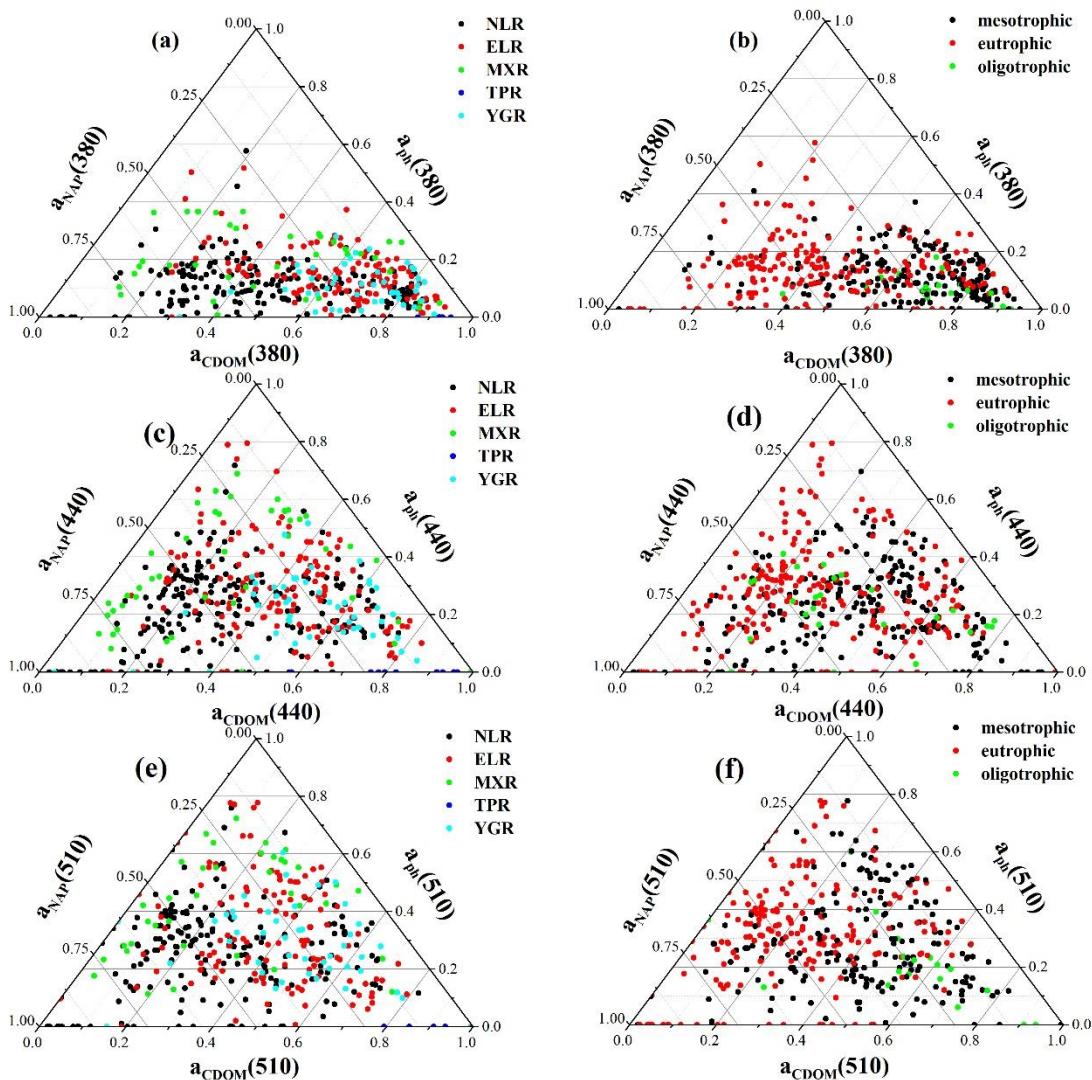
292 Fig. 6 Distribution of CDOM absorption spectral slopes with different wavelength ranges, (a)

293  $S_{CDOM280-400}$ , and (b)  $S_{CDOM275-295}$  in different lake regions.

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### 295 **3.5 Total non-water absorption budget**

296 The absorption of NAP in the NLR and MXR regions generally contributed the most to the total  
297 absorption between 400 and 550 nm while phytoplankton absorption was generally stronger than  
298 CDOM absorption in the MXR region (Fig. S6). For reservoirs in the ELR region, phytoplankton  
299 absorption dominated in the 400-700 nm spectral region. In the YGR region, the contribution of  
300 CDOM absorption was significantly higher than that of other OAC constituents given that these  
301 reservoirs had low eutrophic states. Fig.7 showed that the non-water absorption budget from our  
302 data as triangular diagrams in different hydrological regions and trophic states at 380 nm, 440  
303 nm and 510 nm wavelength. The relative absorption ratios for OACs at 440 nm for various  
304 trophic gradients and hydrological regions was shown in Table S3. And the result of k-means  
305 clustering analysis on non-water absorption budget showed that the dominant water types was  
306 various among regions and trophic states (Table S4 and S5). With increased trophic state, the  
307 relative absorption ratios of phytoplankton absorption and NAP absorption increased while the  
308 ratios of CDOM absorption significantly decreased ( $p < 0.05$ ).



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311 Fig.7 The relative contributions of CDOM, phytoplankton and non-algal particles to total non-water light

312 absorption at 380 nm in different regions(a), and various trophic states(b); at 440 nm in different

313 regions(c), and various trophic states(d); at 510 nm in different regions(e), and various trophic states(f).

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### 316 3.6 Hydrological and water quality parameters for absorption properties

317 The influencing factors for the relative proportion of OACs absorption at 440nm in

318 reservoirs including the hydrological factors and water quality factors were analyzed with  
 319 multiple regression using SPSS 22.0 (Table 3). There are no correlations between absorption  
 320 properties and temperatures ( $p > 0.05$ ) as our sampling fieldtrips were conducted in the autumn  
 321 without obvious seasonal variability. The CDOM proportion was positively related to inactive  
 322 storage (standardized Beta=0.437), DOC (standardized Beta=0.208) and was negatively related  
 323 to TSM (standardized Beta=-0.429) ( $p < 0.005$ ). The TSM concentration (standardized  
 324 Beta=0.472) and total storage (standardized Beta=-0.325) were significantly contributed to NAP  
 325 absorption proportion positively and negatively. The phytoplankton absorption proportion was  
 326 only related to water quality parameters such as *Chla* (standardized Beta=0.802) and TP  
 327 (standardized Beta=0.356).

328 Table 3 The factors affecting OACs absorption proportion with multiple regressions

Types	Factors	Standardized Coefficients (Beta)	t	Significance
<b>CDOM</b>	Constant		13.778	0.000
	Inactive storage	0.437	3.302	0.002
	TSM	-0.429	-3.237	0.003
	DOC	0.208	2.327	0.025
<b>NAP</b>	Constant		12.532	0.000
	TSM	0.472	3.403	0.002
	Total storage	-0.325	-2.344	0.025
<b>Phytoplankton</b>	Constant		2.137	0.038
	<i>Chla</i>	0.802	9.118	0.000
	TP	0.356	2.753	0.010

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## 332 **4 Discussion**

### 333 **4.1 Phytoplankton absorption based on regions and trophic status**

334 The variations of phytoplankton absorption are affected by the size composition variation  
335 of phytoplankton which can often be related to local eutrophic state (Brito et al., 2015; Le et al.,  
336 2013) and geographical settings (Mao et al., 2018), and the climatic conditions could affect the  
337 frequency of algal blooms and the spatial and vertical variations of *Chla*, and the wind-driven  
338 waves affect the bottom re-suspension and change the concentration of OACs (Xue et al., 2017).  
339 The lowest values for phytoplankton absorption coefficients were recorded in the YGR region  
340 mainly due to the low concentration of *Chla* and the relative oligotrophic state of reservoirs in  
341 this region (Astoreca et al., 2012). Conversely, the high concentration of *Chla* in the ELR region  
342 and the high trophic state of reservoirs in the MXR region indicated that reservoirs in these  
343 regions provided a suitable environment for phytoplankton growth, thus accounting for the high  
344 absorption coefficients recorded in reservoirs from these regions (Le et al., 2009). Although  
345 reservoirs in the NLR region were mostly eutrophic, relatively high TSM amounts were  
346 measured probably a consequence of strong spring winds causing the re-suspension of particulate  
347 matter (Li et al., 2015).

348 The variability in  $a^*_{ph}$  within a region indicated a change in pigment composition (Xue et al.,  
349 2017). Comparing mean  $a^*_{ph}$  for the ELR and NLR regions, the higher *Chla* concentration in  
350 ELR and NLR can be attributed to higher concentration of intracellular pigments or cell  
351 aggregation (Shi et al., 2013), and thus could account for the lower  $a^*_{ph}(\lambda)$  values in these  
352 regions. Due to its design (reservoirs divided into regions and trophic state) and its geographical  
353 scope, results of our study should capture variability in geographical conditions and contribute to  
354 improvement in the accuracy of water quality optical models (Shi et al., 2019c), implying a  
355 challenge for parameterization of bio-optical algorithms for different regions (Cao et al., 2015;  
356 Shi et al., 2013).

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## 4.2 Non-algal particle absorption based on regions and trophic status

In the present study, the fact that  $a_{\text{NAP}}^*(440)$  decreased with increasing eutrophication across the five regions (Fig S3) would indicate that the level of eutrophication may affect the distribution of both organic and inorganic particles, and that more inorganic matter is imported into aquatic systems with increased trophic state (Babin et al., 2003). It has further been proposed that a lower  $S_{\text{NAP}}$  corresponds to mineral-dominated waters (Babin et al., 2003). In the present study, the highest  $S_{\text{NAP}}$  values were observed in the YGR region, which was accompanied with the weakest correlation between NAP absorption and MSS. The relative low  $S_{\text{NAP}}$  values in the MXR and NLR regions indicate that the NAP absorption of water reservoirs in these three regions was dominated by mineral matter. This is in good agreement with the strong correlation between NAP and MSS for the MXR and NLR regions. Despite being dominated by inorganic suspended matter, the regional variation in NAP absorption coefficient could probably be due to variation in the composition and size of inorganic particles (Astoreca et al., 2012). The variation in TSM concentration in reservoirs across the five regions could be associated with many factors; geographical characteristics and land-use are important contributing factors to the delivery of terrestrially-derived particles into water reservoirs (Wen et al, 2016). The trophic state of reservoirs could also affect phytoplankton production, microbial activity and degradation process, whereas the delivery of inorganic particles could translate into increase TSM concentration but also the release of particle-bound nutrients which in turn could lead to increased autochthonous production and *Chla* concentration (Astoreca et al., 2012). The complex interrelations between *Chla* and TSM concentration could have variable effects on light absorption. Climatic conditions, such as strong winds, were reported by Wen et al. (2016) to contribute to the re-suspension of

380 bottom sediments in shallow water in the MXR region, resulting in particle absorption.

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### 382 **4.3 CDOM absorptions based on regions and trophic status**

383 The variations of  $S_{280-400}$  for different regions indicated the changes of aromaticity with  
384 humic acids, and fulvic acids (Niu et al., 2014). The  $S_{CDOM}$  values found in this study fell within  
385 the range of previous studies (Song et al., 2010; Niu et al., 2014), yet were higher than results  
386 from Lake Erie and Lake Qinghai (Zhou et al., 2005). The highest mean  $S_{280-400}$  value ( $22.97 \mu\text{m}^{-1}$ )  
387 <sup>1</sup>) was recorded in Yamdrok Lake (TPR). As previously noted by Nima et al. (2016), high altitude,  
388 strong solar irradiance, low trophic states and climatic conditions can decrease phytoplankton  
389 activity, and lead to photo-bleaching and photo-induced degradation organic matters of reservoirs.  
390 In contrast, the lowest mean S value ( $16.9 \mu\text{m}^{-1}$ ) was recorded in eutrophic reservoirs in the YGR  
391 region, indicating a dominance of aromatic materials likely from terrestrial sources (Cory et al.,  
392 2007). Climatic conditions in the YGR region (mean temperature of  $17^\circ\text{C}$  and abundant  
393 precipitation; Liu et al., 2013), were conducive to degradation of plant detritus, the formation of  
394 aromatic humic materials and their transport to aquatic systems in the region (Sun et al., 1997).

395 Also, the negatively significant correlations between  $a^*_{CDOM}(355)$  and  $S_{280-400}$  were observed for all  
396 samples ( $p < 0.05$ ) (Fig.S5) which indicated the composition of CDOM would determine the degree of  
397 DOM color. The higher  $a^*_{CDOM}(355)$  with increased trophic state in our study indicated the production of  
398 fresh DOM with short water retention time and more inputs from the surrounding catchments.

399 Meanwhile, the lower  $a^*_{CDOM}(355)$  and higher S in oligotrophic states and higher suggested more DOM  
400 age and the exposure to photo-chemical degradation process. The differences of correlation coefficients  
401 among various regions indicated the changes of CDOM composition and the degree of photo-degradation  
402 which was due to the content of lignin and the exposure to sunlight (Boyle et al. 2009, Organelli et al.,  
403 2014). With the increasing of trophic states, the inverse relationships indicated the increasing of

404 allochthonous CDOM and the decreasing of photo-degradation process for mesotrophic reservoirs, while  
405 for the eutrophic states, both terrestrial DOM inputs and the strong microbial activities contributed to the  
406 complex trophic states.

#### 407 **4.4 Dominant optical absorptions and regulating factors for management**

408 According to the optical classification of surface waters (based on relative contribution to  
409 the total absorption coefficients of OACs) (Prieur and Sathyendranath, 1981), in terms of trophic  
410 state, the proportion of CDOM absorption decreases while the ratio of NAP absorption to  
411 phytoplankton absorption increases, indicating that CDOM absorption is predominant in  
412 oligotrophic reservoirs while the NAP absorption is predominant in eutrophic waters. Thus, the  
413 reservoirs in the YGR and TPR regions can be classified into the “CDOM-type” whilst the  
414 majority of the reservoirs in the MXR and NLR regions can be classified as the “NAP-type”,  
415 ELR is classified into “phytoplankton-type” according to the relative contribution to the total  
416 absorption coefficients of OACs.

417 The multiple regressions for relative OACs absorption proportions at 440nm based on  
418 hydrological and water quality factors (Table 3) indicated that the hydrological management and  
419 environmental protection for reservoirs would adjust the water optical absorption characteristics  
420 to some extent. The proportion of CDOM in reservoirs was related by inactive storage capacity,  
421 TSM and DOC. This was due to the increasing inactive storage capacity would extend the  
422 hydraulic retention time, and the decreasing of TSM concentration would relatively decrease the  
423 proportion of NAP absorption and increase the relative proportion of CDOM and phytoplankton  
424 absorption. The increasing of total storage would introduce lots of sediments and terrestrial input  
425 of solid matter which contributed to the accumulation of non-algal particles and the increasing  
426 proportion of NAP absorption. For phytoplankton absorption proportion, it was only closed to  
427 the water quality parameters including *Chla* and TP rather than other hydrological conditions.



428 Our results demonstrated the variation of the dominant absorption agent in different hydrological  
429 regions and trophic states and offered the detailed regulation conditions with hydrological and  
430 water chemical parameters to control the optical absorption characteristics (Shi et al., 2013). Also,  
431 the specific absorption coefficient for a given OAC should be different when bio-optical models  
432 are developed for remotely monitoring water quality constituents of reservoirs across China.

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## 434 **5 Conclusions**

435 This study represents a systematic investigation of the variability of absorption properties of  
436 OACs in inland waters. The variation of sources and composition of OACs would affect the  
437 correlations between optical absorption parameters and relative water chemical parameters which  
438 is primary for remote monitoring of water quality in reservoirs for further study. It provides  
439 unique insights into the factors controlling that variability and requires for the parameterization  
440 of optical models in different regions. In addition, the differences of optical absorption  
441 characteristics in various regions which will help us to establish the more accurate remote  
442 sensing models of quasianalytical algorithm in different regions with remote sensing data for  
443 next step. The well understanding of spatial variations of IOPs are needed to meet the challenges  
444 to the variations of regional bio-optical modelling and water-quality parameter remote sensing  
445 algorithm parameterization and performance. Moreover, the good and significant relationships  
446 between the optical absorption characteristics and water quality parameters would provide more  
447 potential ways to calibrate the water quality parameters' monitoring models indirectly in future  
448 studies. It provides a simple outline for the classification of optically-complex waters, and  
449 therefore contribute to a new stage in the development of water type-specific algorithms.

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