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# Tephrostratigraphy and tephrochronology of lakes Ohrid and Prespa, Balkans

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## Abstract

Four cores from Balkans lakes Ohrid and Prespa were studied for recognition of tephra layers and cryptotephra, and the results presented along with the review of data from other two already published cores from Lake Ohrid. The six cores provide a previously unrealised tephrostratigraphic framework of the two lakes, and supply the first detailed tephrochronologic profile (composite) for the Balkans, which spans from the end of the Middle Pleistocene to the end of the Ancient Age (AD 472). A total of 12 tephra layers and cryptotephra were recognised in the cores. One is of Middle Pleistocene age (131 ky) and correlated to the marine tephra layer P-11 from Pantelleria Island. Eight volcanic layers are Upper Pleistocene in age, and encompass the period between ca. 107 ky and ca. 31 ky. This interval contains some of the main regional volcanic markers of the Central Mediterranean area, including X-6, X-5, Y-5 and Y-3 tephra layers. The other layers of this interval have been related to the marine tephra layers C20, Y-6 and C10, while one was for the first time recognised in distal areas and correlated to the Taurano eruption of probable Vesuvian origin. Three cryptotephra were of Holocene age. Two of which have been correlated to Mercato and AD 472 eruptions of Somma-Vesuvius, while the third has been correlated to the FL eruption from Mount Etna. These recognitions provide a link of the Ohrid and Prespa lacustrine successions to other archives of the Central Mediterranean area, like South Adriatic, Ionian, and South Tyrrhenian Seas, lakes of Southern Italy (Lago Grande di Monticchio, Pantano di San Gregorio Magno and Lago di Pergusa) and Balkans (Lake Shkodra).

## 1 Introduction

Owing to the intense explosive volcanic activity that affected the Mediterranean over the last 200 ky (Vezzoli, 1988; Poli et al., 1987; Santacroce, 1987; Rosi and Sbrana, 1987; Keller et al., 1990; Orsi et al., 1996; Pappalardo et al., 1999; Di Vito et al., 2008; Santacroce et al., 2008) application of tephrochronology to volcanology, Quaternary

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science, paleoceanography, and archaeology has an exceedingly high potential in this area. In the last 30 y, Quaternary tephra layers have been extensively used to develop a high-resolution event stratigraphy for the late Pleistocene and Holocene across the Central and Eastern Mediterranean (Keller et al., 1978; Paterne et al., 1988, 1990; Narcisi and Vezzoli, 1999; Wulf et al., 2004; Margari et al., 2007; Aksu et al., 2008; Giaccio et al., 2009; Zanchetta et al., 2010). While early work focused mainly on samples from marine cores (Keller et al., 1978; Paterne et al., 1988, 1990; Calanchi et al., 1998), recent work on terrestrial archives (including Italian, Greek, Turkish and Balkan lakes, and Bulgarian, Greeks and Italian cave sites; e.g., St Seymour and Christianis, 1995; Narcisi and Vezzoli, 1999; St Seymour et al., 2004; Wulf et al., 2004; Frisia et al., 2008; Giaccio et al., 2008, 2009; Sulpizio et al., 2010; Vogel et al., 2010) have considerably advanced the development of a long, high-resolution tephrostratigraphy, which will link marine and terrestrial records of Pleistocene-Holocene-age across the Mediterranean region and mainland Europe. Nevertheless, there are some areas in the Central Mediterranean, in which volcanologic and Quaternary studies are still at the beginning, and the link to the general tephrostratigraphic Mediterranean network is lacking. The Balkans area across Macedonia, Albania and Montenegro is certainly one, in which tephrostratigraphic and tephrochronologic studies are in their infancy, although some studies on lacustrine settings indicate the area is extremely promising for tephrostratigraphic studies (Wagner et al., 2008; Caron et al., 2010; Lézine et al., 2010; Sulpizio et al., 2010; Vogel et al., 2010). For these areas both tephrostratigraphic and tephrochronologic studies can offer invaluable stratigraphic support for sedimentologic, palaeoclimatic, paleoenvironmental, and volcanological studies. Here we present tephrostratigraphic and tephrochronologic data of lakes Ohrid and Prespa (Fig. 1), which encompass the last 131 ky.

The paper reviews some already published data from the Macedonian side of Lake Ohrid, which are presented along with new data from both Ohrid and Prespa lakes. The recognised tephra layers supply a composite tephrostratigraphy of lakes Ohrid and Prespa, which has been correlated to other archives from the Balkans (Albanian side of

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lake Ohrid, Caron et al., 2010; Lézine et al., 2010; Lake Shkodra, Sulpizio et al., 2010), and linked to the regional tephrostratigraphic network of the Central Mediterranean area.

## 2 Study sites and analytical methods

Lake Ohrid (Fig. 1) is a transboundary lake shared by the Republics of Albania and Macedonia. It is located at 693 m a.s.l., and surrounded by high mountain ranges reaching heights up to 2300 m. It is 30 km long, 15 km wide and covers an area of 358 km<sup>2</sup>. The lake basin shows a relatively simple tub-shaped morphology with a maximum water depth of 289 m, an average water depth of 151 m and a total volume of 50.7 km<sup>3</sup> (Popovska and Bonacci, 2007).

Lake Ohrid is mainly fed by inflow from karst springs (55%), while the remaining 45% of the hydrological input includes direct precipitation on the lake surface, river and direct surface runoff (Matzinger et al., 2006). The direct watershed of Lake Ohrid covers an area of 1002 km<sup>2</sup> (Popovska and Bonacci, 2007). Surface outflow (60%) through the River Crn Drim to the north and evaporation (40%) are the main hydrological outputs (Matzinger et al., 2006). The average annual precipitation on the Lake Ohrid watershed is 907 mm (Popovska and Bonacci, 2007).

Lake Prespa is located 10 km to the east of Lake Ohrid, at an altitude of 849 m a.s.l. It is a transboundary lake shared between the Republics of Macedonia, Albania, and Greece. The surface area is 254 km<sup>2</sup>, with a catchment area of ca. 1300 km<sup>2</sup>, a maximum water depth of 48 m, and a volume of 3.6 km<sup>3</sup>. The total inflow is estimated to 16.9 m<sup>3</sup> s<sup>-1</sup>, with 56% originating from river runoff from numerous small streams, 35% from direct precipitation, and 9% from Lake Mikri Prespa to the south (Matzinger et al., 2006). Water loss derives through evaporation (52%), irrigation (2%) and outflow through karst aquifers (46%).

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## 2.1 Core recovery and description

The sediment cores presented here were recovered during field campaigns between 2005 and 2009, using a floating platform equipped with a gravity corer for surface sediments and a piston corer (both UWITEC Co.) for deeper sediments. Cores Lz1120, Co1200, Co1201, and Co1202 are from Lake Ohrid and cores Co1204 and Co1216 from lake Prespa (Fig. 1). Core description and high-resolution colour scanning was carried out immediately after lengthwise opening of the cores in the laboratory.

Figure 2 summarises the major lithological features of all cores presented in this study. Detailed lithological descriptions are presented elsewhere (Lz1120, Wagner et al., 2009; Co1200 and Co1201, Lindhorst et al., 2010; Co1202, Vogel et al., 2010a, b; Co1204, Wagner et al., 2010).

## 2.2 Tephra detection and analysis

High-resolution XRF analysis was carried out on the surface of one of the core halves using an ITRAX core scanner (COX Ltd), equipped with a Mo-tube set to 30 kV and 30 mA, and a Si-drift chamber detector. Scanning was performed at 0.5 mm (Co1204), 1 mm (Co1202), 2 mm (Co1216), and 2.5 mm (Co1200, Co1201) resolution and an analysis time of 20 s (Co1200, Co1201, Co1202, Co1204) and 10 s (Co1216) per measurement. The obtained count rates for K, Sr and Zr can be used as estimates of the relative concentrations for these elements.

From horizons indicating conspicuousness in macroscopic grain size composition, colour or element count rates derived from XRF scanning, about 1 cm<sup>3</sup> was washed and sieved. The >40 mm fraction was embedded in epoxy resin and screened for glass shards and micro-pumice fragments using scanning electron microscopy (SEM). Energy-dispersive spectrometry (EDS) of glass shards and micro-pumice fragments was performed using an EDAX-DX micro-analyser mounted on a Philips SEM 515 (operating conditions: 20 kV acceleration voltage, 100 s live time counting, 200–500 nm beam diameter, 2100–2400 shots s<sup>-1</sup>, ZAF correction). The ZAF correction procedure

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does not include natural or synthetic standards for reference, and requires analysis normalisation at a given value (chosen at 100%). Analytical precision is 0.5% for abundances higher than 15 wt%, 1% for abundances around 5 wt%, 5% for abundances of 1 wt%, and less than 20% for abundances close to the detection limit (around 0.5 wt%).

5 Interlaboratory standards are shown in Table 1.

Accuracy of measurements is around 1%, a value analogous to that obtained using wave dispersion spectroscopy (WDS), as tested by Cioni et al. (1998) and Marianelli and Sbrana (1998). Comparison of EDS and WDS micro-analyses carried out on the same samples has shown differences less than 1% for abundances greater than 10 0.5 wt% (e.g., Cioni et al., 1998; Sulpizio et al., 2010) confirming the full comparability of EDS analyses from the Pisa laboratory and data from WDS microprobes.

The concentration of thirty-five trace elements was determined by ICP-MS for four selected samples. About 50–60 mg of sample powders were dissolved in screw-top PFA vessels with a mixture of HF and HNO<sub>3</sub> on a hot-plate at ~120 °C. The sample 15 solutions were then spiked with Rh, Re and Bi as internal standards (20 ng ml<sup>-1</sup> in the final solutions) and diluted to 50 ml in polypropylene flasks. Milli-Q purified water (18.2 MΩ cm), and HF and HNO<sub>3</sub> Aristar grade were used in each step of sample preparation. Analyses were performed by external calibration using geochemical reference samples as composition- and matrix-matching calibration solutions. The correction procedure includes (i) blank subtraction; (ii) instrumental drift correction using 20 internal standardization and repeated (every 5 samples) analysis of a drift monitor; (iii) oxide-hydroxide interference correction. At the concentration levels of the studied samples, precisions are better than 5% RSD, except for Sc, Ni and Cu for which the precisions are between 5 and 10% RSD.

### 25 3 Description of tephra layers

The studied sediment cores contain a variable number of tephra and cryptotephra, which are here described on the basis of their lithology and composition. Tephra and

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cryptotephra numbering follows the origin of the core (OT for Ohrid Tephra and PT for Prespa Tephra), the year of core recovery (05-09) and the last two identification numbers of the respective core. Glass shards and micropumice fragments from the different cores were compositionally classified using the silica vs. total alkali diagram (TAS; Le Bas et al., 1986; Fig. 3).

### 3.1 Core Co1200

Core Co1200 contains two discrete tephra layers visible at naked eye. Tephra layer OT0700-1 (40–38 cm) appears light-brown to ochre in colour, and contains highly vesicular, aphyric micro-pumice and glass shards (Fig. 4a). Most of the glass shards show a trachytic composition with only few samples that plot close to the phonolitic field on the TAS diagram (Fig. 3a).

Tephra layer OT0700-2 (120.5–85.5 cm) has a light red colour, and contains aphyric, vesicular micro-pumices and aphyric glass shards with thick septa (Fig. 4b). Glass composition ranges from trachyte to phono-trachyte (Fig. 3a), with three different alkali ratios (Table 2).

### 3.2 Core Co1201

Core Co1201 contains seven discrete tephra layers visible at naked eye inspection and with different thicknesses.

The five tephra layers between 81–78 cm (OT0701-1), 90–89 cm (OT0701-2), at 100 cm (OT0701-3), and between 106–104 cm (OT0701-4) and 126–110 cm (OT0701-5) are light red to rusty-red in colour and comprise aphyric micro-pumice and glass shards. They show identical glass composition (Fig. 3b), and are grouped into three different compositional groups on the basis of the different alkali ratios (Table 2).

Tephra layer OT0701-6 (186–184 cm), appears pink in colour, and contains vesicular, aphyric micro-pumice fragments and glass shards (Fig. 4c). The glass composition is homogeneous phonolitic (Fig. 3b) with alkali ratio greater than 1.5 (Table 2).

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Tephra layer OT0701-7 (192–190 cm), is light-brown to red in colour, and contains both aphyric glass shards and poorly vesicular fragments with small pyroxene crystals in the groundmass (Fig. 4d). The glass composition continuously spans between phono-tephryte/shoshonite and phonolite/trachyte (Fig. 3b). Three main compositional groups can be identified on the basis of the alkali ratio and SiO<sub>2</sub> content (Table 2).

### 3.3 Core Co1202

Core Co1202 contains four visible tephra layers and six cryptotephra, which were already described in detail by Vogel et al. (2010a).

Cryptotephra OT0702-1 (77.5–74.5 cm) comprises dark-brown tachylitic fragments with crystalline groundmass (mainly sanidine, clinopyroxene and leucite; Fig. 4e). The composition straddles the foiditic and tephri-phonolitic fields (Fig. 3c).

Cryptotephra OT0702-2 (145.5–144 cm) consists of non-vesicular and blocky fragments with a porphyritic texture. Fragments exhibit mineral inclusions of plagioclase, clinopyroxene and olivine up to some tens of microns in size (Fig. 4f) and frequent occurrences of Fe-Ti oxides. The glass composition is mostly benmoreitic, with few analyses plotting in the trachytic and mugearitic fields on the TAS diagram (Fig. 3c).

Cryptotephra OT0702-3 (277.5–269 cm) comprises mainly aphyric, vesicular micropumice. The bubbles of the micropumices are mainly circular with thick septa (Fig. 4g). The glass composition of the cryptotephra is a fairly homogeneous Na-phonolite (Table 2 and Fig. 3c).

Tephra layer OT0702-4 (620–617 cm) is light brown in colour, and is characterised by normally graded middle to fine sand. Glass shards are mixed with lacustrine sediment up to 2 cm above the tephra layer. The tephra comprises highly vesicular, aphyric micro-pumice and cusped glass shards. Most of the glass shards show a trachytic composition with only few samples that plot close to the phonolitic field on the TAS diagram (Fig. 3c). Trace element distribution (sample 523; Table 3, Fig. 5) shows the less enriched pattern in the four analysed samples, with small negative anomalies in Ba

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and Sr and a marked enrichment in Pb (Fig. 5a). The rare earth element (REE) pattern shows a more or less regular decrease passing from light REE (LREE) to heavy REE (HREE; Fig. 5b).

Cryptotephra OT0702-5 (696–689 cm) comprises tachilitic particles with a crystal-rich groundmass containing acicular clinopyroxene, plagioclase and sanidine (Fig. 4h). The few glass compositions available range from latite to phonolite when plotted on the TAS diagram (Table 2 and Fig. 3c).

Tephra layer OT0702-6 (752–743 cm) is reddish-brown to light-brown in colour, and glass shards are mixed with lacustrine sediments in the overlying 10 cm. Volcanic particles mainly comprise aphyric, vesicular micro-pumices and aphyric glass shards with thick septa. Glass composition ranges from trachyte to phono-trachyte (Fig. 3c), with two different alkali ratio (Table 2). Trace element distribution (sample 565; Table 3, Fig. 5) shows an intermediate enrichment in the analysed samples, with a pronounced negative anomalies in Sr and Eu, moderate anomaly in Ba, and moderate positive anomalies in Th, U and Pb (Fig. 5a). The REE pattern shows a regular decrease in enrichment passing from LREE to HREE, with a pronounced negative Eu anomaly (Fig. 5b).

Cryptotephra OT0702-7 (825–822 cm) comprises aphyric, cusped glass shards with thin septa and glassy groundmass. Glass composition ranges from rhyolitic (main) to trachytic (Table 2 and Fig. 3c).

Tephra layer OT0702-8 (1146.5–1140 cm) is rusty-red in colour and comprises coarse ash and shows sharp basal and top contacts. Volcanic particles are aphyric micro-pumices and glass shards. Glass composition is mainly phonolitic (Table 2 and Fig. 3c).

Tephra layer OT0702-9 (1232.5–1229 cm) is light-brown in colour, and shows sharp basal and top boundaries. Grain size mainly comprises fine to coarse ash. Volcanic particles are highly vesicular, aphyric micro-pumices and glass shards with a large variability in shape and size (Fig. 4i). Glass composition is mainly trachytic, with few analyses that plot into the phonolitic field, and two different alkali ratios (Table 2 and

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Fig. 3c). The trace element (sample 1003; Fig. 5, Table 3) distribution shows the most enriched pattern in the four analysed samples, with marked negative anomalies in Ba, Sr and Eu that testify for a feldspar-dominated magma fractionation (Fig. 5a). The REE pattern shows a regular decrease from LREE to HREE, with a marked negative Eu anomaly (Fig. 5b).

Cryptotephra OT0702-10 (1447–1440 cm) comprises mainly aphyric cusped glass shards. When plotted on the TAS diagram the glass shards reveal a bimodal chemical composition, which comprises a trachyte and a rhyolite without any compositional trend between (Fig. 3c).

### 3.4 Core Lz1120

Core Lz1120 contains two visible tephra layers and one cryptotephra, which were already described in detail by Wagner et al. (2008).

Cryptotephra OT0520-1 (315–310 cm) comprises dark brown, blocky fragments with few spherical or ovoid vesicles. The groundmass comprises small crystals of plagioclase and minor olivine. Single-shard analyses show dispersion between mugearitic and benmoreitic fields (Fig. 3d).

Tephra layer OT0520-2 (897–896 cm) comprises light coloured, elongated, highly vesicular fragments. Vesicles are mainly tubular and form channels throughout the entire length of the pyroclastic fragments. Groundmass is glassy, with very few elongated microcrystals of sanidine. Composition of single shards shows a limited variability within the trachytic field (Fig. 3d).

Tephra OT0520-3 (1075–1070 cm) is of unknown thickness, as the resistance of this layer prevented penetration of the coring equipment through it. Volcanic fragments comprise mainly light coloured, highly vesicular fragments, with minor dark coloured glass shards and micropumice fragments. In both typologies of volcanic fragments the vesicles are spherical or ovoidal, separated by thin, glassy sets. Groundmass is almost aphyric, even if small sanidine crystals sometimes occur on larger sets among

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bubbles. Single shard composition show a trend from the trachytic to the phonolitic fields (Fig. 3d), and can be arranged into three groups on the basis of different alkali ratios (Table 2).

### 3.5 Core Co1205

5 Core Co1204 was recovered from the north-western part of Lake Prespa (Fig. 1). It contains two discrete tephra layers and one cryptotephra.

Tephra layer PT0704-1 (672.5–667.5 cm) has a light grey colour, and contains aphyric, highly vesicular micro-pumice fragments and glass shards with glassy groundmass. The glass composition is homogeneously trachytic (Fig. 3e; Table 2).

10 Cryptotephra PT0704-2 (767.2–764.2 cm) was identified by high Sr count rates through high-resolution XRF-scanning, and contains tachilitic fragments with highly crystalline groundmass. Glass is rare and shows a compositional trend from shoshonites to trachytes (Fig. 3e). Three different compositional groups were identified on the basis of SiO<sub>2</sub>, CaO, and total alkali contents (Table 2).

15 Tephra layer PT0704-3 (879.3–863.3 cm) contains aphyric, highly vesicular micro-pumice and cusped to convolute glass shards with glassy groundmass. Glass composition straddles the phonolitic and trachytic fields (Fig. 3e), it can be split into two or three groups depending on the different alkali ratio (Table 2). The trace element distribution and the REE pattern (sample PR628; Table 3, Fig. 5) are identical to the  
20 sample 565 from core Co1202.

## 4 Discussion

### 4.1 Correlation to proximal deposits and other distal archives

The correlation of a distal tephra layer with proximal counterparts is a critical process, which in many cases implies the contemporaneous use of different data, such as glass

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and mineral composition, chronology, lithology, and stratigraphic position. This is because pyroclastic deposits from the same source, with few exceptions, show striking major element composition. Furthermore, pyroclastic deposits from different sources but originating from magmas with similar degree of evolution (e.g. trachytes and rhyolites) are hardly distinguishable on the basis of the sole major element composition. In this study, cores Lz1120 and Co1202, which contain correlated tephra layers (Wagner et al., 2008; Vogel et al., 2010), can be used as a reference, particularly because core Co1202 contains the largest number of tephra layers in all studied sediment cores from this region (Fig. 2).

The youngest volcanic deposit was correlated to the AD 472 (1478 cal. y BP) eruption of Somma-Vesuvius. It occurs as a cryptotephra in core Co1202 (OT0702-1), but has not been recognised in any of the other studied cores (Fig. 2).

The FL ( $3370 \pm 70$  cal. y BP; Coltelli et al., 2000) cryptotephra occurs in cores Lz1120 (OT0520-1; Fig. 2; Wagner et al., 2008) and Co1202 (OT0702-2; Fig. 2; Vogel et al., 2010) from Lake Ohrid, but has not been found in the successions from lake Prespa.

The Mercato ( $8540 \pm 50$  cal. y BP; Zanchetta et al., 2010) cryptotephra occurrence is limited to core Co1202 (OT0702-3; Fig. 2; Vogel et al., 2010) from Lake Ohrid.

Due to its thickness and unique, homogeneous trachytic composition (Table 2) the Y-3 tephra layer was recognized in cores Lz1120 (OT0520-2; Wagner et al., 2008), Co1200 (OT0700-1), and Co1202 (OT0702-4; Vogel et al., 2010) from Lake Ohrid and in core Co1204 (PT0704-1) from Lake Prespa (Fig. 2; Table 2). The ICP-MS analyses of trace elements on sample OT0702-4 (Table 3) reinforces the correlation to the proximal deposits of the SMP1-e eruption ( $30.67 \pm 0.23$  cal. ky BP) from the Campi Flegrei caldera, which is indicated as the proximal counterpart of the Y-3 tephra layer (Di Vito et al., 2008; Zanchetta et al., 2008). The occurrence of the Y-3 tephra layer in four out of six cores (Fig. 2) indicates its usefulness as a stratigraphic marker for lakes Ohrid and Prespa.

The Codola ash (inferred age 33 cal. ky BP; Giaccio et al., 2008 or 34.2 cal. ky BP; Vogel et al., 2010) occurs as a cryptotephra in core Co1202 (OT0702-5; Fig. 2;

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Vogel et al., 2010). The Codola fragments have a highly microcrystalline groundmass and a glass composition that straddles the tephri-phonolitic/latitic and the phonolitic/trachytic (Fig. 6a; Di Vito et al., 2008; Giaccio et al., 2008; Santacroce et al., 2008). Based on their stratigraphic position and lithology the tephra layers PT0704-2 in core Co1204 and PT0916-1 in core Co1216 from Lake Prespa are good candidates for correlation to the Codola eruption. The inspection of glass composition illustrates a more complex situation, with some of the analyses from the PT0704-2 and the PT0916-1 samples that plot outside the Codola field (Fig. 6a). In particular, the PT0916-1 analyses define two compositional groups separated by a broad gap in SiO<sub>2</sub> and total alkali content (PT0916-1a and PT0916-1b; Fig. 6a; Table 2). Both groups plot outside the Codola field, either in TAS or SiO<sub>2</sub> vs. CaO diagrams (Fig. 5a, b), being more and less evolved, respectively. The less evolved analyses correlate well to the Taurano composition, which defines an evolutionary trend with the Codola samples in the SiO<sub>2</sub> vs. CaO diagram (Fig. 6b). The Taurano eruption was tentatively assigned to the activity of the Somma-Vesuvius volcano, and approximately dated 36–33 cal. ky BP (Di Vito et al., 2008), however this age range was subsequently changed to 30–33 cal. ky BP on the basis of the correlation to distal tephra layers in the Lago Grande di Monticchio succession (Sulpizio et al., 2008). Proximal deposits comprise porphyritic dark scoriae of shoshonitic/phono-tephritic composition, with a groundmass rich in microlites of clinopyroxene, sanidine and leucite (Di Vito et al., 2008). The evolved group (sample PT0916-1b; Table 2) plots outside the Taurano-Codola trend, and coincides with the Campanian Ignimbrite field in SiO<sub>2</sub> vs. CaO diagram (pinkish area of Fig. 5b). The lithology of the glass shards supports the correlation of this group to the Campanian Ignimbrite deposits, being formed by convolute and cusped shards with glassy groundmass. The occurrence of mixed populations of Taurano and Campanian Ignimbrite ash fragments is not surprising, due to the stratigraphic vicinity of the two eruptions (Di Vito et al., 2008). The analyses of cryptotephra PT0704-2 from core Co1204 can be divided into three groups, on the basis of different contents of SiO<sub>2</sub>, total alkali, and CaO (Fig. 5a, b; Table 2). Among them, only the PT0704-2b samples plot in the Codola

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field in both TAS and SiO<sub>2</sub> vs. CaO diagram (Fig. 5a, b). The two analyses of group PT0704-2c are the most evolved, and plot in the Campanian Ignimbrite field in the SiO<sub>2</sub> vs. CaO diagram (Fig. 6b). The PT0704-2a group have shoshonitic/latitic composition (Fig. 6a), and shows a separate compositional trend in the SiO<sub>2</sub> vs. CaO diagram. The correlation of this group with the regional tephrostratigraphy is at present not possible, since tephra layers or cryptotephra with similar lithology, glass composition and age have never been described in the Balkans, in Adriatic/Ionian marine cores, and continental Italy (e.g., Lago Grande di Monticchio succession; Wulf et al., 2004). Based on glass composition, a generic correlation to the activity of Vulcano Island between 53 and 21 ky (De Astis et al., 2000; 2006; Lucchi et al., 2008) is here proposed, although the link to a specific dated eruption is, to date, impossible.

The Campanian Ignimbrite/Y-5 (CI/Y-5; 39.28±0.11 ky; De Vivo et al., 2001) tephra layer has previously been recognized in cores Lz1120 (OT0520-3; Fig. 2; Wagner et al., 2008) and Co1202 (OT0702-6; Fig. 2; Vogel et al., 2010a) from Lake Ohrid. The glass composition of the OT0520-3 and the OT0702-6 tephra layers straddles the trachytic and phonolitic fields (Fig. 3c), and can be organised into two to three different compositional groups on the basis of alkali ratios (Table 2). These compositional groups reflect the involvement in the eruptive processes of three differently evolved trachytic magmas (i.e. from “evolved” to “primitive” trachytes; Fedele et al., 2008), which provides a geochemical fingerprint of the CI/Y-5 deposits also in distal reaches (Civetta et al., 1997; Pappalardo et al., 2002). Based on their lithology, stratigraphic position and major element glass composition the tephra layers found in cores Co1200 (OT0700-2; Fig. 2), Co1201 (OT0701-1, -2, -3, -4, -5; Fig. 2) from Lake Ohrid and in core Co1204 (PT0704-3; Fig. 2) from Lake Prespa can also be correlated to OT0520-3 and OT0702-6 and thus to the CI/Y-5 eruption. ICP-MS analyses of trace elements of samples OT0702-6 and PT0704-3 (Table 3; Fig. 5) reinforce this correlation. If we also consider the glass shards mixed into the PT0916-1 cryptotephra, the CI/Y-5 distal ash occurs in all studied cores from lakes Ohrid and Prespa (Fig. 2). The CI/Y-5 tephra layers is the lowermost volcanic deposit in most of the studied cores, with the only exceptions

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of cores Co1201 and Co1202 (Fig. 2).

Cryptotephra OT0702-7 occurs below the CI/Y-5 tephra layer in core Co1202, and was correlated to the Green Tuff eruption from Pantelleria island (Vogel et al., 2010), which corresponds to the Y-6 marine tephra layer (Keller et al., 1978).

5 The tephra OT0701-6 occurs below the CI/Y-5 in core Co 1201, and has a homogeneous trachytic composition (Fig. 3b; Table 2). It has no correspondence in the other cores from lakes Ohrid and Prespa, but can be correlated to the C20 marine tephra layer (dated at 79–80 ky; Paterne et al., 1988), and to the proximal deposits of SA3-b eruption from the Campanian area (Table 3).

10 The tephra layer OT0702-8 occurs in core Co1202 between 825 cm and 822 cm, and was correlated to the X-5 marine tephra layer ( $105 \pm 2$  ky; Keller et al., 1978; Kraml, 1997), which has been tentatively related to the TAU1-b eruption from Campi Flegrei caldera (Di Vito et al., 2008). Due to its chemical composition the OT0701-3 cryptotephra from core Co1201 can tentatively be correlated to the OT0702-8 tephra layer  
15 from core Co1202 and thus to the TAU1-b/X-5 eruption.

Tephra layer OT0702-9 in core Co1202 was correlated to the marine X-6 tephra layer (Keller et al., 1978), of generic Campanian origin. Brauer et al. (2007) quoted an  $^{40}\text{Ar}/^{39}\text{Ar}$  age of  $107 \pm 2$  ky for the X-6 tephra, which is in good agreement with the suggested age of 108.43 ky obtained from the varve-supported chronology of the Lago  
20 Grande di Monticchio record. The X-6 tephra layer has not been recognised in the other studied cores of lakes Ohrid and Prespa (Fig. 2). The evolved trace element composition of this tephra layer (Fig. 5; Table 3) can help in its discrimination from the mess of trachytic tephra layers recognised in the Upper Pleistocene successions.

25 Cryptotephra OT0702-10 (Fig. 2; Vogel et al., 2010a) in core Co1202 from Lake Ohrid represents the deepest volcanic deposit recognised in cores from lakes Ohrid and Prespa. It correlates to the marine P11 tephra layer (ca. 131 ka; Paterne et al., 1988, 2008), which sourced from Pantelleria Island.

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## 4.2 Composite tephrochronological record and regional correlations

The correlation of the 12 recognised tephra and cryptotephra layers with known proximal deposits allows the reconstruction of a composite stratigraphic and chronological framework for lakes Ohrid and Prespa during the past 131 ky (Fig. 7). Three cryptotephra occur during Holocene, with the Mercato layer marking the temperature maximum of the Early Holocene. The following seven tephra layers and cryptotephra punctuate the Upper Pleistocene, encompassing about 77 ky (from 30 to 107 ky; Fig. 7). Four markers cluster between 30 and 40 ky (i.e. between the Y-3 and the Y-5 tephra layers), detailing the stratigraphic succession between Heinrich events 3 and 4 (Ton-That et al., 2001; Zanchetta et al., 2008; Wagner et al., 2010). Particularly significant is the first recognition of the Taurano deposits in a very distal succession, which so far has never been described in nearby Adriatic and Ionian Sea cores. The Y-6 cryptotephra occurs below the CI, at about 49 ky (Fig. 7), linking the Balkans to the Ionian Sea archives. About 30 ky of sedimentation occur below the Y-6 without any tephra record, being the following tephra layer correlated to the SA3-a eruption and the C-20 marine tephra layer (79–80 ky; Paterne et al., 1988; Fig. 7). The last two tephra layers of the Upper Pleistocene are the X-5 and X-6, dated at 105 and 107 ky, respectively (Fig. 7). The oldest and deepest volcanic deposit recognised is the P-11 tephra layer, which anticipates the inception of the Last Interglacial in the Balkans (Lézine et al., 2010; Vogel et al., 2010b).

The recognised tephra layers and cryptotephra can be correlated to other published tephrostratigraphic records of the Central Mediterranean area, linking the Balkans with other archives at a regional scale (Fig. 8). Excluding the CI/Y-5 tephra layer, which extensively occurs in the Central Mediterranean area and extends through Aegean Sea and up to Russia (Pyle et al., 2006; Giaccio et al., 2008), the other recognised volcanic markers have various frequency of occurrence in the published archives (Fig. 8). The most common is the Y-3 tephra layer, which commonly occur in marine cores from South Adriatic (SA), Ionian, and South Tyrrhenian (ST) Seas, and in Lago Grande di

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Monticchio succession (Fig. 8). Other widespread layers are the Mercato, X-5, and X-6, while the AD 472, and FL deposits link the Balkans to the Monticchio succession, to the Campanian area, and to the Sicily (Fig. 8). The P-11 cryptotephra is the only, together with the Y-5 tephra layer, that links the Balkans to the Ionian Sea and to the Aegean Sea. This is because the pantelleritic/trachytic deposit recognised in a core from Lesvos Island, and attributed to the Pantelleria Green Tuff/Y6 (Margari et al., 2007), was reinterpreted as the P-11 deposit on the basis of geochemical data (Vogel et al., 2010).

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## Tephra layers in lakes Ohrid and Prespa

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**Table 1.** Comparison of the EDS device used in this study in comparison with WDS micro-probes from GeoForschungsZentrum (GFZ, Potsdam, Germany), and from Saclay (France; Cioni et al., 1998). Reference material comprises volcanic glasses with a chemistry ranging from basalt to trachyte. ALV981R23, CFA47, and KE12 samples from Cioni et al. (1998).

	ALV981R23 (basalt) EDS n=12				CFA 47 (trachyte) EDS n=12				KE 12 (pantellerite) EDS n=10				SMP1-a (trachyte) EDS n=13				CA1-a (trachyte) EDS n=18			
	sd	Saclay	sd	sd	sd	Saclay	sd	sd	sd	Saclay	sd	sd	sd	GFZ	sd	sd	sd	GFZ	sd	
SiO <sub>2</sub>	49.56	0.14	49.79	0.19	61.39	0.27	61.94	0.33	70.8	0.23	70.83	0.22	60.29	0.16	60.17	0.17	58.41	0.45	58.67	0.39
TiO <sub>2</sub>	1.30	0.07	1.28	0.05	0.48	0.09	0.42	0.05	0.31	0.06	0.28	0.02	0.48	0.07	0.46	0.03	0.53	0.09	0.48	0.01
Al <sub>2</sub> O <sub>3</sub>	16.57	0.13	16.67	0.08	18.61	0.11	18.62	0.16	7.92	0.16	7.82	0.03	19.23	0.1	19.4	0.09	19.24	0.18	18.96	0.07
FeO	8.44	0.22	8.46	0.09	2.76	0.12	2.66	0.15	8.41	0.06	8.67	0.2	3.01	0.08	3.15	0.05	3.95	0.24	4.22	0.14
MnO	0.22	0.03	0.14	0.06	0.24	0.13	0.18	0.04	0.38	0.07	0.29	0.03	0.29	0.13	0.26	0.04	0.18	0.09	0.14	0.03
MgO	8.82	0.16	8.73	0.11	0.58	0.08	0.42	0.02	0.1	0.04	0	0	0.22	0.06	0.36	0.02	0.99	0.15	1.01	0.1
CaO	11.84	0.25	11.87	0.11	1.83	0.08	1.85	0.06	0.36	0.07	0.35	0.01	1.67	0.06	1.87	0.05	3.22	0.25	3.5	0.23
Na <sub>2</sub> O	3.00	0.18	2.9	0.04	5.43	0.08	5.4	0.12	7.1	0.19	7.23	0.25	6.65	0.23	6.34	0.17	4.28	0.2	4.27	0.1
K <sub>2</sub> O	0.1	0.05	0.05	0.01	8.14	0.04	8.02	0.16	4.29	0.07	4.19	0.15	7.23	0.21	7.1	0.08	8.6	0.29	8.21	0.26
P <sub>2</sub> O <sub>5</sub>	–	–	–	–	–	–	–	–	–	–	–	–	0	0	0	0	0	0	0	0
Cl	–	–	–	–	–	–	–	–	–	–	–	–	0.92	0.05	0.9	0.02	0.6	0.06	0.55	0.03
S	0.14	0.04	0.12	0	0.54	0.08	0.49	0.01	0.32	0.08	0.33	0	–	–	–	–	–	–	–	–

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**Table 2.** Average EDs analyses of glass shards of tephra layers and cryptotephra from the studied and reviewed cores.

Sample	depth (cm)	analyses	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO <sub>tot</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	ClO	Total	Total alk	Alk. Ratio	
Co1200	OT0700-1	40–38	<i>n</i> =12 <b>sd</b>	61.6 <b>0.79</b>	0.38 <b>0.1</b>	18.74 <b>0.16</b>	3.03 <b>0.35</b>	0.1 <b>0.09</b>	0.67 <b>0.16</b>	2.3 <b>0.31</b>	3.74 <b>0.45</b>	8.94 <b>0</b>	0.5 <b>0.14</b>	100	12.68	2.39	
	OT0700-2a	120.5–85.5	<i>n</i> =11 <b>sd</b>	60.8 <b>0.29</b>	0.37 <b>0.08</b>	19.22 <b>0.14</b>	2.96 <b>0.08</b>	0.18 <b>0.06</b>	0.39 <b>0.07</b>	1.7 <b>0.08</b>	6.44 <b>0.18</b>	7.24 <b>0.1</b>	0 <b>0.07</b>	100	13.69	1.12	
	OT0700-2b	120.5–85.5	<i>n</i> =4 <b>sd</b>	61.17 <b>0.16</b>	0.44 <b>0.13</b>	19.32 <b>0.13</b>	2.89 <b>0.07</b>	0.3 <b>0.04</b>	0.4 <b>0.05</b>	1.73 <b>0.07</b>	5.74 <b>0.09</b>	7.33 <b>0.16</b>	0 <b>0.03</b>	100	13.08	1.28	
	OT0700-2c	120.5–85.5	<i>n</i> =1 <b>sd</b>	59.85 <b>0.16</b>	0.45 <b>0.13</b>	18.73 <b>0.13</b>	3.64 <b>0.04</b>	0.12 <b>0.07</b>	0.84 <b>0.05</b>	2.71 <b>0.07</b>	3.18 <b>0.16</b>	10.19 <b>0</b>	0 <b>0.03</b>	100.01	13.37	3.2	
Co1201	OT0701-1/5a	126–78	<i>n</i> =36 <b>sd</b>	60.79 <b>0.3</b>	0.41 <b>0.09</b>	19.19 <b>0.19</b>	2.98 <b>0.12</b>	0.23 <b>0.09</b>	0.41 <b>0.09</b>	1.67 <b>0.09</b>	6.4 <b>0.21</b>	7.22 <b>0.2</b>	0 <b>0</b>	100	13.62	1.13	
	OT0701-1/5b	126–78	<i>n</i> =11 <b>sd</b>	61.32 <b>0.37</b>	0.4 <b>0.12</b>	19.13 <b>0.33</b>	2.99 <b>0.17</b>	0.14 <b>0.12</b>	0.48 <b>0.13</b>	1.87 <b>0.26</b>	5.53 <b>0.43</b>	7.55 <b>0.4</b>	0 <b>0</b>	100	13.08	1.37	
	OT0701-1/5c	126–78	<i>n</i> =9 <b>sd</b>	60.75 <b>0.59</b>	0.31 <b>0.1</b>	18.97 <b>0.56</b>	3.27 <b>0.25</b>	0.06 <b>0.06</b>	0.75 <b>0.11</b>	2.45 <b>0.17</b>	3.36 <b>0.41</b>	9.76 <b>0.31</b>	0 <b>0</b>	100	13.12	2.91	
	OT0701-6	186–184	<i>n</i> =15 <b>sd</b>	59.58 <b>0.23</b>	0.44 <b>0.07</b>	20.04 <b>0.12</b>	2.76 <b>0.14</b>	0.13 <b>0.09</b>	0.51 <b>0.09</b>	2.31 <b>0.09</b>	5.24 <b>0.16</b>	8.8 <b>0.17</b>	0 <b>0</b>	100	14.04	1.68	
	OT0701-7a	192–190	<i>n</i> =8 <b>sd</b>	50.86 <b>1.79</b>	1.15 <b>0.16</b>	18.55 <b>0.33</b>	8.33 <b>1.26</b>	0.22 <b>0.06</b>	3.19 <b>0.59</b>	8.47 <b>1.15</b>	3.27 <b>0.45</b>	5.4 <b>0.78</b>	0.28 <b>0.14</b>	0.29 <b>0.04</b>	99.99	8.66	1.65
	OT0701-7b	192–190	<i>n</i> =8 <b>sd</b>	58.3 <b>1.48</b>	0.63 <b>0.12</b>	19.75 <b>0.2</b>	3.62 <b>0.89</b>	0.15 <b>0.08</b>	0.82 <b>0.43</b>	3.41 <b>1.07</b>	4.54 <b>0.31</b>	8.36 <b>0.31</b>	0.01 <b>0.62</b>	0.4 <b>0.05</b>	99.99	12.9	1.84
	OT0701-7c	192–190	<i>n</i> =1 <b>sd</b>	62.55 <b>0.16</b>	0.6 <b>0.12</b>	18.22 <b>0.2</b>	2.75 <b>0.08</b>	0.18 <b>0.08</b>	0.49 <b>0.43</b>	0.97 <b>1.07</b>	7.26 <b>0.31</b>	6.43 <b>0.31</b>	0 <b>0.62</b>	0.56 <b>0.05</b>	100.01	13.69	0.89
Co1202	OT0702-1	77.5–74.5	<i>n</i> =7 <b>sd</b>	<b>48.82</b> <b>1.12</b>	<b>0.88</b> <b>0.14</b>	<b>20.58</b> <b>0.84</b>	<b>7.08</b> <b>0.58</b>	<b>0.27</b> <b>0.1</b>	<b>1.39</b> <b>0.34</b>	<b>8.04</b> <b>1.51</b>	<b>6.19</b> <b>0.58</b>	<b>5.61</b> <b>0.72</b>	<b>0.06</b> <b>0.04</b>	<b>1.08</b> <b>0.1</b>	100	11.8	0.92
	OT0702-2	145.5–144	<i>n</i> =10 <b>sd</b>	55.98 <b>1.92</b>	1.62 <b>0.26</b>	18.18 <b>1.26</b>	6.52 <b>1.88</b>	0.21 <b>0.13</b>	1.94 <b>0.77</b>	5.35 <b>0.95</b>	5.86 <b>0.49</b>	3.63 <b>0.69</b>	0.4 <b>0.11</b>	0.3 <b>0.09</b>	100	9.49	0.62
	OT0702-3	277.5–269	<i>n</i> =9 <b>sd</b>	59.1 <b>0.55</b>	0.17 <b>0.09</b>	21.58 <b>0.13</b>	1.95 <b>0.11</b>	0.17 <b>0.07</b>	1.16 <b>0.09</b>	1.76 <b>0.13</b>	7.56 <b>0.56</b>	7.02 <b>0.27</b>	0 <b>0</b>	0.52 <b>0.03</b>	100	14.58	0.93
	OT0702-4	620–617	<i>n</i> =12 <b>sd</b>	61.27 <b>0.64</b>	0.39 <b>0.06</b>	18.7 <b>0.12</b>	3.11 <b>0.21</b>	0.11 <b>0.07</b>	0.67 <b>0.18</b>	2.35 <b>0.25</b>	3.73 <b>0.46</b>	9.22 <b>0.48</b>	0 <b>0</b>	0.46 <b>0.1</b>	100	12.96	2.52
	OT0702-5	696–689	<i>n</i> =4 <b>sd</b>	57.54 <b>0.34</b>	0.79 <b>0.18</b>	20.54 <b>1.02</b>	3.7 <b>0.73</b>	0.17 <b>0.09</b>	0.61 <b>0.15</b>	5.16 <b>0.62</b>	3.58 <b>0.5</b>	7.58 <b>0.62</b>	0 <b>0</b>	0.35 <b>0.11</b>	100.02	11.16	2.12
	OT0702-6a	752–743	<i>n</i> =10 <b>sd</b>	60.71 <b>0.23</b>	0.42 <b>0.07</b>	19.17 <b>0.08</b>	2.95 <b>0.13</b>	0.21 <b>0.09</b>	0.45 <b>0.06</b>	1.7 <b>0.11</b>	6.49 <b>0.15</b>	7.22 <b>0.14</b>	0 <b>0</b>	0.69 <b>0.04</b>	100	13.71	1.11
	OT0702-6b	752–743	<i>n</i> =2 <b>sd</b>	<b>61.27</b> <b>0.23</b>	<b>0.31</b> <b>0.24</b>	<b>19.01</b> <b>0.21</b>	<b>2.89</b> <b>0.18</b>	<b>0.13</b> <b>0.18</b>	<b>0.58</b> <b>0.16</b>	<b>2.05</b> <b>0.36</b>	<b>5.5</b> <b>0.59</b>	<b>7.81</b> <b>0</b>	<b>0</b> <b>0</b>	<b>0.47</b> <b>0.14</b>	100	13.31	1.42
	OT0702-7a	825–822	<i>n</i> =10 <b>sd</b>	71.89 <b>0.47</b>	0.45 <b>0.1</b>	8.31 <b>1.25</b>	8.07 <b>0.18</b>	0.34 <b>0.08</b>	0.09 <b>0.1</b>	0.34 <b>0.16</b>	5.24 <b>0.71</b>	4.36 <b>0.16</b>	0 <b>0</b>	0.92 <b>0.16</b>	100	9.65	0.85
	OT0702-7b	825–822	<i>n</i> =2 <b>sd</b>	68.04 <b>1.64</b>	0.59 <b>0.11</b>	11.37 <b>1.25</b>	7.76 <b>0.18</b>	0.35 <b>0.08</b>	0.27 <b>0.1</b>	0.71 <b>0.16</b>	5.72 <b>0.71</b>	4.69 <b>0.16</b>	0 <b>0</b>	0.52 <b>0.16</b>	100	10.41	0.83
	OT0702-8	1146.5–1140	<i>n</i> =12 <b>sd</b>	57.6 <b>0.48</b>	0.55 <b>0.08</b>	19.49 <b>0.16</b>	4.48 <b>0.34</b>	0.16 <b>0.07</b>	1.22 <b>0.16</b>	3.77 <b>0.33</b>	4.18 <b>0.19</b>	8.09 <b>0.45</b>	0.01 <b>0.03</b>	0.46 <b>0.05</b>	100	12.26	1.94
	OT0702-9	1232.5–1229	<i>n</i> =15 <b>sd</b>	61.15 <b>0.31</b>	0.46 <b>0.07</b>	18.82 <b>0.12</b>	3.09 <b>0.14</b>	0.29 <b>0.09</b>	0.39 <b>0.11</b>	1.68 <b>0.08</b>	6.39 <b>0.64</b>	7.03 <b>0.52</b>	0 <b>0</b>	0.71 <b>0.15</b>	100	13.42	1.12
	OT0702-10a	1447–1440	<i>n</i> =3 <b>sd</b>	66.18 <b>0.89</b>	0.43 <b>0.16</b>	16.23 <b>3.11</b>	4.14 <b>3.16</b>	0.23 <b>0.2</b>	0.2 <b>0.21</b>	0.84 <b>0.12</b>	5.85 <b>0.79</b>	5.79 <b>1.39</b>	0 <b>0</b>	0.13 <b>0.11</b>	100.01	11.64	0.99
	OT0702-10b	1447–1440	<i>n</i> =6 <b>sd</b>	72.52 <b>0.79</b>	0.41 <b>0.09</b>	9.34 <b>1.23</b>	6.7 <b>0.4</b>	0.31 <b>0.06</b>	0.12 <b>0.06</b>	0.34 <b>0.07</b>	5.06 <b>0.71</b>	4.5 <b>0.31</b>	0 <b>0</b>	0.71 <b>0.08</b>	100	9.55	0.91

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Table 2. Continued.

Sample	depth (cm)	analyses	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO <sub>tot</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	ClO	Total	Total alk	Alk. Ratio		
Lz1120	OT0520-1	315–310	<i>n</i> =15	54.25	1.76	17.48	8.15	0	2.85	6.02	5.4	3.29	0.48	0.31	100	8.69	0.61	
			<b>sd</b>	<b>0.73</b>	<b>0.22</b>	<b>0.51</b>	<b>0.53</b>	<b>0</b>	<b>0.33</b>	<b>0.5</b>	<b>0.41</b>	<b>0.34</b>	<b>0.1</b>	<b>0.06</b>				
	OT0520-2	897–896	<i>n</i> =12	61.88	0.29	18.63	2.92	0.03	0.5	2.2	4.18	8.82	0	0.54	99.99	12.99	2.11	
			<b>sd</b>	<b>0.52</b>	<b>0.08</b>	<b>0.19</b>	<b>0.23</b>	<b>0.06</b>	<b>0.15</b>	<b>0.19</b>	<b>0.47</b>	<b>0.51</b>	<b>0</b>	<b>0.14</b>				
	OT0520-3a	1075–1070	<i>n</i> =3	61.16	0.43	18.96	3.03	0.23	0.35	1.72	6.17	7.08	0.05	0.81	100	13.25	1.15	
			<b>sd</b>	<b>0.15</b>	<b>0.02</b>	<b>0.07</b>	<b>0.09</b>	<b>0.01</b>	<b>0</b>	<b>0.05</b>	<b>0.13</b>	<b>0.35</b>	<b>0.01</b>	<b>0.01</b>				
OT0520-3b	1075–1070	<i>n</i> =6	61.94	0.43	19.09	3.03	0.2	0.46	1.97	4.79	7.39	0.07	0.65	100	12.18	1.54		
		<b>sd</b>	<b>0.38</b>	<b>0.04</b>	<b>0.2</b>	<b>0.07</b>	<b>0.05</b>	<b>0.13</b>	<b>0.25</b>	<b>0.43</b>	<b>0.56</b>	<b>0.04</b>	<b>0.19</b>					
OT0520-3c	1075–1070	<i>n</i> =1	61.29	0.42	19.08	3.15	0.11	0.61	2.41	3.85	8.38	0.19	0.51	100	12.24	2.18		
		<b>sd</b>	–	–	–	–	–	–	–	–	–	–	–					
Co1204	PT0704-1	672.5–667.5	<i>n</i> =10	61.43	0.37	18.62	3.17	0.1	0.66	2.34	3.54	9.34	0	0.43	100.01	12.88	2.64	
			<b>sd</b>	<b>0.59</b>	<b>0.06</b>	<b>0.11</b>	<b>0.24</b>	<b>0.07</b>	<b>0.15</b>	<b>0.22</b>	<b>0.41</b>	<b>0.39</b>	<b>0</b>	<b>0.11</b>				
	PT0704-2a	767.2–764.2	<i>n</i> =4	54.54	0.59	22.5	3.46	0.07	1.52	9.79	3.03	4.22	0.03	0.25	99.99	7.25	1.39	
			<b>sd</b>	<b>1.63</b>	<b>0.39</b>	<b>1.92</b>	<b>1.43</b>	<b>0.07</b>	<b>0.94</b>	<b>1.54</b>	<b>0.75</b>	<b>0.86</b>	<b>0.06</b>	<b>0.14</b>				
	PT0704-2b	767.2–764.2	<i>n</i> =5	57.66	0.68	22.18	2.85	0	0.54	5.29	3.37	7.14	0.07	0.22	100	10.5	2.12	
			<b>sd</b>	<b>0.63</b>	<b>0.11</b>	<b>1.85</b>	<b>1.22</b>	<b>0</b>	<b>0.31</b>	<b>0.66</b>	<b>0.28</b>	<b>0.75</b>	<b>0.11</b>	<b>0.11</b>				
	PT0704-2c	767.2–764.2	<i>n</i> =2	61.45	0.41	18.99	3.06	0.11	0.69	2.18	4.05	8.59	0	0.48	100	12.64	2.12	
			<b>sd</b>	<b>0.65</b>	<b>0</b>	<b>0.07</b>	<b>0.35</b>	<b>0.16</b>	<b>0.35</b>	<b>0.62</b>	<b>1.48</b>	<b>0.98</b>	<b>0</b>	<b>0.07</b>				
	PT0704-3a	879.3–863.3	<i>n</i> =25	60.77	0.44	19.13	2.93	0.26	0.4	1.64	6.38	7.33	0	0.71	100	13.72	1.15	
			<b>sd</b>	<b>0.24</b>	<b>0.08</b>	<b>0.13</b>	<b>0.08</b>	<b>0.07</b>	<b>0.07</b>	<b>0.09</b>	<b>0.19</b>	<b>0.19</b>	<b>0</b>	<b>0.05</b>				
PT0704-3b	879.3–863.3	<i>n</i> =10	61.34	0.43	19.45	3.06	0.17	0.59	1.85	5.13	7.47	0	0.5	99.99	12.6	1.46		
		<b>sd</b>	<b>0.9</b>	<b>0.08</b>	<b>0.43</b>	<b>0.29</b>	<b>0.08</b>	<b>0.25</b>	<b>0.48</b>	<b>0.57</b>	<b>0.71</b>	<b>0</b>	<b>0.14</b>					
PT0704-3c	879.3–863.3	<i>n</i> =4	61.22	0.41	19.08	3.27	0.11	0.79	2.2	3.26	9.31	0	0.36	100	12.57	2.86		
		<b>sd</b>	<b>1.02</b>	<b>0.08</b>	<b>0.17</b>	<b>0.29</b>	<b>0.13</b>	<b>0.17</b>	<b>0.36</b>	<b>0.46</b>	<b>0.93</b>	<b>0</b>	<b>0.08</b>					
Co1216	PT0916-1a	428.6–425.8	<i>n</i> =6	59.26	0.49	20.21	2.8	0.17	0.72	2.54	4.96	8.45	0	0.42	100.01	13.41	1.7	
			<b>sd</b>	<b>0.36</b>	<b>0.06</b>	<b>0.19</b>	<b>0.18</b>	<b>0.04</b>	<b>0.07</b>	<b>0.15</b>	<b>0.25</b>	<b>0.19</b>	<b>0</b>	<b>0.06</b>				
			<i>n</i> =6	50.22	1.09	19.05	8.12	0.17	3.68	9.39	2.91	4.78	0.33	0.26	99.99	7.69	1.64	
			<b>sd</b>	<b>0.7</b>	<b>0.14</b>	<b>0.42</b>	<b>0.49</b>	<b>0.09</b>	<b>0.33</b>	<b>0.74</b>	<b>0.29</b>	<b>0.49</b>	<b>0.11</b>	<b>0.06</b>				

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**Table 3.** ICP-MS trace element analyses of four selected samples from lakes Ohrid and Prespa: 523, 565, and 1033 are from core Co1202 (Lake Ohrid), PR628 from core Co1204 (Lake Prespa). Some analyses are reported for comparison and correlation: 896-Oh from core Lz1120 (Di Vito et al., 2008; Zanchetta et al., 2008); JO2004Y3 from core JO2004 (Albanian side of Lake Ohrid; Caron et al., 2010; Lézine et al., 2010); TM15 from Lago Grande di Monticchio succession (Wulf et al., 2004; sample courtesy of S. Wulf); SMP1-e from proximal deposits of SMP1-e/Y-3 deposits (Di Vito et al., 2008); JO2004Y5 from core JO2004 (Albanian side of Lake Ohrid; Caron et al., 2010; Lézine et al., 2010).

T. layer sample	Y3					Y5			X6
	OT0702-4 523	896-Oh	JO2004Y3	TM15	SMP1-e	OT0702-6 565	PT0704-3 PR628	JO2004Y5	OT0702-9 1033
Li	37	40	33	36	49	64	59	64	73
Be	8.1	8	7.7	8.6	12.2	16.2	17.3	17	15.4
Sc	7	7	6	4	4	6	5	5	6
V	59	70	47	52	37	41	35	23	40
Cr	15	51	7	1	1	22	3	20	23
Co	8	11	9	3	2	5	3	3	8
Ni	21	58	47	4	1	24	3	28	33
Cu	7	11	11	4	4	10	6	14	7
Ga	16.2	17.5	15.3	16.4	18	20.9	19.7	19.7	19.6
Rb	254	246	255	286	321	335	337	343	299
Sr	412	361	464	534	176	128	157	129	113
Y	30.1	29.5	27.4	27.8	36	54	52	52	66
Zr	266	238	252	264	373	539	541	543	622
Nb	37	34	35	36	51	87	86	88	97
Cs	15.7	14.6	15	15.6	21.3	29.5	29.5	29.1	28.4
Ba	1058	962	898	735	76	858	644	243	288
La	59	55	55	55	69	107	108	106	150
Ce	113	106	107	105	134	207	208	197	297
Pr	12.7	12	11.7	11.7	14.6	22.3	22.3	21.5	31.7
Nd	45	42	41	42	51	75	75	71	106
Sm	8.2	7.8	7.6	7.5	9.2	12.9	12.9	12.6	18.1
Eu	1.95	1.79	1.87	1.88	1.69	1.65	1.64	1.35	1.74
Gd	6.4	6.2	6.2	6	6.9	10.5	10.3	10	14.2
Tb	0.94	0.97	0.88	0.88	1.08	1.58	1.55	1.55	2.1
Dy	5	5.1	4.6	4.8	5.8	8.8	8.4	8.3	11.2
Ho	1.01	1	0.93	0.91	1.15	1.79	1.71	1.69	2.22
Er	2.63	2.66	2.47	2.44	3.09	4.9	4.7	4.6	5.8
Tm	0.41	0.41	0.35	0.38	0.49	0.73	0.71	0.7	0.89
Yb	2.46	2.58	2.26	2.41	3.11	4.6	4.5	4.5	5.7
Lu	0.36	0.36	0.33	0.34	0.44	0.67	0.65	0.68	0.82
Hf	6.1	5.6	5.8	6	8.2	11.6	11.7	11.5	13.3
Ta	2.13	2	2.02	2.01	2.83	4.7	4.6	4.8	4.9
Tl	2.1	2.1	2	2	2.35	3	3	2.4	2.4
Pb	43	39	41	42	51	54	53	57	55
Th	22.9	21	21.9	21.9	30.7	42	42	41	51
U	6.5	6	6.2	6.7	8.8	13.5	13.3	13.7	14.7

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**Table 4.** Selected analyses (mainly EDS; WDS, EPMA and only for one case XRF) for tephra layers correlated with those found in core Co1202. Data from Lago Grande di Monticchio are from Wulf et al. (2004). Data from S. Gregorio Magno are from Munno and Petrosino (2007). Data for proximal deposits of Mercato are from Santacroce et al. (2008). For Y-3 data see Zanchetta et al. (2008) and Wagner et al. (2008). For Codola see Di Vito et al. (2008), Giaccio et al. (2008). For CI/Y-5 see Di Vito et al. (2008), Giaccio et al. (2008), and Wagner et al. (2008). For the Y-6 see Margari et al. (2007). For P-11 see Paterne et al. (2008).

			SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO <sub>tot</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	ClO	Total	Total alkali	alkali ratio	
AD 472	OT0702-1	<i>n</i> =7	48.82	0.88	20.58	7.08	0.27	1.39	8.04	6.19	5.61	0.06	1.08	100	11.8	0.91	
		<b>sd</b>	<b>1.12</b>	<b>0.14</b>	<b>0.84</b>	<b>0.58</b>	<b>0.1</b>	<b>0.34</b>	<b>1.51</b>	<b>0.58</b>	<b>0.72</b>	<b>0.04</b>	<b>0.1</b>				
		<i>n</i> =333	49.79	0.5	22.23	4.95	0.15	1.01	5.63	9.2	5.41	0	1.12	100	14.61	0.59	
	Proximal	<b>sd</b>	<b>1.65</b>	<b>0.16</b>	<b>0.92</b>	<b>1.38</b>	<b>0.07</b>	<b>0.47</b>	<b>1.51</b>	<b>2.35</b>	<b>2.29</b>	<b>0</b>	<b>0.53</b>				
		<i>n</i> =28	50.28	0.55	22.64	4.91	0.17	0.65	5.69	8.32	5.82	0.02	0.97	100.02	14.14	0.7	
		<b>sd</b>	<b>0.83</b>	<b>0.11</b>	<b>0.62</b>	<b>0.53</b>	<b>0.07</b>	<b>0.29</b>	<b>0.67</b>	<b>1.02</b>	<b>0.92</b>	<b>0.03</b>	<b>0.1</b>				
	LGM/TM2b	<i>n</i> =6	50.18	0.5	21.61	4.79	0.2	0.64	5.89	7.9	6.92	0.29	1.07	99.99	14.82	0.88	
		<b>sd</b>	<b>1.02</b>	<b>0.07</b>	<b>0.19</b>	<b>0.24</b>	<b>0.04</b>	<b>0.2</b>	<b>0.61</b>	<b>0.35</b>	<b>0.72</b>	<b>0.17</b>	<b>0.19</b>				
FL	OT0702-2	<i>n</i> =10	55.98	1.62	18.18	6.52	0.21	1.94	5.35	5.86	3.63	0.4	0.3	99.99	9.49	0.62	
		<b>sd</b>	<b>1.92</b>	<b>0.26</b>	<b>1.26</b>	<b>1.88</b>	<b>0.13</b>	<b>0.77</b>	<b>0.95</b>	<b>0.49</b>	<b>0.69</b>	<b>0.11</b>	<b>0.09</b>				
		<i>n</i> =15	54.25	1.76	17.48	8.15	0	2.85	6.02	5.4	3.29	0.48	0.31	99.99	8.69	0.61	
	OT0520-1	<b>sd</b>	<b>0.73</b>	<b>0.22</b>	<b>0.51</b>	<b>0.53</b>	<b>0</b>	<b>0.33</b>	<b>0.5</b>	<b>0.41</b>	<b>0.34</b>	<b>0.1</b>	<b>0.06</b>				
		<i>n</i> =14	56.81	1.21	20.52	4.86	0.09	1.38	5.27	6.22	3.1	0.24	0.3	100	9.32	0.5	
		<b>sd</b>	<b>1.78</b>	<b>0.44</b>	<b>2.39</b>	<b>2.49</b>	<b>0.09</b>	<b>1.27</b>	<b>1.83</b>	<b>1.04</b>	<b>1.26</b>	<b>0.22</b>	<b>0.11</b>				
	Shkodra	<i>n</i> =10	53.02	1.9	17.27	8.93	0.22	2.98	5.94	5.09	3.95	0.43	0.27	100	9.04	0.78	
		<b>sd</b>	<b>1.03</b>	<b>0.5</b>	<b>1.5</b>	<b>1.95</b>	<b>0.1</b>	<b>0.89</b>	<b>1.18</b>	<b>0.5</b>	<b>1.43</b>	<b>0.15</b>	<b>0.12</b>				
Mercato	OT0702-3	<i>n</i> =9	59.1	0.17	21.58	1.95	0.17	0.16	1.76	7.56	7.02	0	0.52	99.99	14.58	0.93	
		<b>sd</b>	<b>0.55</b>	<b>0.09</b>	<b>0.13</b>	<b>0.11</b>	<b>0.07</b>	<b>0.09</b>	<b>0.13</b>	<b>0.56</b>	<b>0.27</b>	<b>0</b>	<b>0.03</b>				
		<i>n</i> =40	58.51	0.13	21.7	1.76	0.14	0.09	1.66	8.56	6.93	0	0.52	100	15.49	0.81	
	Proximal	<b>sd</b>	<b>0.72</b>	<b>0.08</b>	<b>0.41</b>	<b>0.19</b>	<b>0.1</b>	<b>0.08</b>	<b>0.26</b>	<b>0.62</b>	<b>0.38</b>	<b>0</b>	<b>0.09</b>				
		<i>n</i> =10	58.68	0.14	21.41	1.8	0.18	0.07	1.76	8.58	6.76	0.02	0.58	99.98	15.34	0.79	
		<b>sd</b>	<b>0.3</b>	<b>0.03</b>	<b>0.21</b>	<b>0.13</b>	<b>0.03</b>	<b>0.01</b>	<b>0.25</b>	<b>0.18</b>	<b>0.54</b>	<b>0.02</b>	<b>0.04</b>				
SMP1-e/Y-3	OT0700-1	<i>n</i> =12	61.6	0.38	18.74	3.03	0.1	0.67	2.3	3.74	8.94	0	0.5	100	12.68	2.39	
		<b>sd</b>	<b>0.79</b>	<b>0.1</b>	<b>0.16</b>	<b>0.35</b>	<b>0.09</b>	<b>0.16</b>	<b>0.31</b>	<b>0.45</b>	<b>0.45</b>	<b>0</b>	<b>0.14</b>				
		<i>n</i> =12	61.27	0.39	18.7	3.11	0.11	0.67	2.35	3.73	9.22	0	0.46	100.01	12.95	2.47	
	OT0702-4	<b>sd</b>	<b>0.64</b>	<b>0.06</b>	<b>0.12</b>	<b>0.21</b>	<b>0.07</b>	<b>0.18</b>	<b>0.25</b>	<b>0.46</b>	<b>0.48</b>	<b>0</b>	<b>0.1</b>				
		<i>n</i> =10	61.43	0.37	18.62	3.17	0.1	0.66	2.34	3.54	9.34	0	0.43	100.01	12.88	2.64	
		<b>sd</b>	<b>0.59</b>	<b>0.06</b>	<b>0.11</b>	<b>0.24</b>	<b>0.07</b>	<b>0.15</b>	<b>0.22</b>	<b>0.41</b>	<b>0.39</b>	<b>0</b>	<b>0.11</b>				
	Proximal	<i>n</i> =10	62.42	0.51	18.38	2.87	0.24	0.47	1.92	4.79	8.41	n.a.	n.a.	100.01	13.2	1.76	
		<b>sd</b>	<b>0.18</b>	<b>0.06</b>	<b>0.08</b>	<b>0.08</b>	<b>0.08</b>	<b>0.08</b>	<b>0.06</b>	<b>0.15</b>	<b>0.13</b>	<b>0</b>	<b>0</b>				
		<i>n</i> =19	62.22	0.38	18.36	3.27	0.13	0.61	2.19	3.85	8.36	0.12	0.52	100.01	12.21	2.17	
	LGM/TM15	<b>sd</b>	<b>0.78</b>	<b>0.03</b>	<b>0.21</b>	<b>0.29</b>	<b>0.04</b>	<b>0.15</b>	<b>0.22</b>	<b>0.44</b>	<b>0.55</b>	<b>0.06</b>	<b>0.11</b>				
		<i>n</i> =10	61.79	0.4	18.57	3.29	0.14	0.64	2.52	3.69	8.96	0	0	100	12.65	2.43	
		<b>sd</b>	<b>1.21</b>	<b>0.05</b>	<b>0.22</b>	<b>0.32</b>	<b>0.09</b>	<b>0.19</b>	<b>0.4</b>	<b>0.44</b>	<b>0.95</b>	<b>0</b>	<b>0</b>				
Taurano	PT0916-1a	<i>n</i> =6	50.22	1.09	19.05	8.12	0.17	3.68	9.39	2.91	4.78	0.33	0.26	99.99	7.69	1.64	
		<b>sd</b>	<b>0.7</b>	<b>0.14</b>	<b>0.42</b>	<b>0.49</b>	<b>0.09</b>	<b>0.33</b>	<b>0.74</b>	<b>0.29</b>	<b>0.49</b>	<b>0.11</b>	<b>0.06</b>				
		<i>n</i> =72	50.7	0.84	19.34	7.04	0.17	2.37	8.49	3.43	6.5	0.55	0.58	100	9.93	1.9	
	Proximal	<b>sd</b>	<b>1.7</b>	<b>0.12</b>	<b>0.94</b>	<b>0.81</b>	<b>0.04</b>	<b>0.95</b>	<b>1.14</b>	<b>0.85</b>	<b>0.89</b>	<b>0.3</b>	<b>0.15</b>				

Table 4. Continued.

		SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO <sub>tot</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	ClO	Total	Total alkali	alkali ratio		
Codola	OT0702-5	<i>n</i> =4	57.54	0.79	20.54	3.7	0.17	0.61	5.16	3.58	7.58	0	0.35	100.02	11.16	2.12	
		<b>sd</b>	<b>0.34</b>	<b>0.18</b>	<b>1.02</b>	<b>0.73</b>	<b>0.09</b>	<b>0.15</b>	<b>0.62</b>	<b>0.5</b>	<b>0.62</b>	<b>0</b>	<b>0.11</b>				
	PT0704-2b	<i>n</i> =5	57.66	0.68	22.18	2.85	0	0.54	5.29	3.37	7.14	0.07	0.22	100	10.5	2.12	
		<b>sd</b>	<b>0.63</b>	<b>0.11</b>	<b>1.85</b>	<b>1.22</b>	<b>0</b>	<b>0.31</b>	<b>0.66</b>	<b>0.28</b>	<b>0.75</b>	<b>0.11</b>	<b>0.11</b>				
	Proximal-a	<i>n</i> =10	53.18	0.78	18.6	7.12	0.17	1.87	6.85	3.09	7.31	0.47	0.58	100.02	10.4	2.37	
		<b>sd</b>	<b>1.04</b>	<b>0.04</b>	<b>0.82</b>	<b>0.51</b>	<b>0.03</b>	<b>0.56</b>	<b>1.54</b>	<b>0.29</b>	<b>0.63</b>	<b>0.04</b>	<b>0.18</b>				
	Proximal-b	<i>n</i> =10	57.92	0.53	19.74	3.56	0.15	0.59	4.28	3.57	9.06	0.11	0.51	100.02	12.63	2.54	
		<b>sd</b>	<b>0.07</b>	<b>0.03</b>	<b>0.13</b>	<b>0.17</b>	<b>0.01</b>	<b>0.05</b>	<b>0.06</b>	<b>0.05</b>	<b>0.14</b>	<b>0.02</b>	<b>0.05</b>				
	LGM/TM16a	<i>n</i> =11	52.39	0.79	19.04	7.06	0.16	1.85	7.13	3.65	6.66	0.71	0.56	100	10.31	1.82	
		<b>sd</b>	<b>1.69</b>	<b>0.09</b>	<b>0.86</b>	<b>0.6</b>	<b>0.03</b>	<b>0.31</b>	<b>1.24</b>	<b>0.43</b>	<b>0.62</b>	<b>1.05</b>	<b>0.11</b>				
	LGM/TM16b	<i>n</i> =6	58.58	0.5	19.7	3.66	0.12	0.55	4.3	3.4	8.58	0.1	0.5	99.99	11.98	2.52	
		<b>sd</b>	<b>0.65</b>	<b>0.06</b>	<b>0.41</b>	<b>0.33</b>	<b>0.03</b>	<b>0.14</b>	<b>0.95</b>	<b>0.49</b>	<b>0.4</b>	<b>0.04</b>	<b>0.08</b>				
CI/Y-5	OT0700-2a	<i>n</i> =11	60.8	0.37	19.22	2.96	0.18	0.39	1.7	6.44	7.24	0	0.7	100	13.69	1.12	
		<b>sd</b>	<b>0.29</b>	<b>0.08</b>	<b>0.14</b>	<b>0.08</b>	<b>0.06</b>	<b>0.07</b>	<b>0.08</b>	<b>0.18</b>	<b>0.1</b>	<b>0</b>	<b>0.07</b>				
	OT0700-2b	<i>n</i> =4	61.17	0.44	19.32	2.89	0.3	0.4	1.73	5.74	7.33	0	0.69	100	13.08	1.28	
		<b>sd</b>	<b>0.16</b>	<b>0.13</b>	<b>0.13</b>	<b>0.07</b>	<b>0.04</b>	<b>0.05</b>	<b>0.07</b>	<b>0.09</b>	<b>0.16</b>	<b>0</b>	<b>0.03</b>				
	OT0701-1/5a	<i>n</i> =36	60.79	0.41	19.19	2.98	0.23	0.41	1.67	6.4	7.22	0	0.7	100	13.62	1.13	
		<b>sd</b>	<b>0.3</b>	<b>0.09</b>	<b>0.19</b>	<b>0.12</b>	<b>0.09</b>	<b>0.09</b>	<b>0.09</b>	<b>0.21</b>	<b>0.2</b>	<b>0</b>	<b>0.05</b>				
	OT0701-1/5b	<i>n</i> =11	61.32	0.4	19.13	2.99	0.14	0.48	1.87	5.53	7.55	0	0.58	100	13.08	1.37	
		<b>sd</b>	<b>0.37</b>	<b>0.12</b>	<b>0.33</b>	<b>0.17</b>	<b>0.12</b>	<b>0.13</b>	<b>0.26</b>	<b>0.43</b>	<b>0.4</b>	<b>0</b>	<b>0.18</b>				
	OT0701-1/5c	<i>n</i> =9	60.75	0.31	18.97	3.27	0.06	0.75	2.45	3.36	9.76	0	0.32	100	13.12	2.91	
		<b>sd</b>	<b>0.59</b>	<b>0.1</b>	<b>0.56</b>	<b>0.25</b>	<b>0.06</b>	<b>0.11</b>	<b>0.17</b>	<b>0.41</b>	<b>0.31</b>	<b>0</b>	<b>0.07</b>				
	OT0702-6a	<i>n</i> =10	60.71	0.42	19.17	2.95	0.21	0.45	1.7	6.49	7.22	0	0.69	100	13.71	1.11	
		<b>sd</b>	<b>0.23</b>	<b>0.07</b>	<b>0.08</b>	<b>0.13</b>	<b>0.09</b>	<b>0.06</b>	<b>0.11</b>	<b>0.15</b>	<b>0.14</b>	<b>0</b>	<b>0.04</b>				
	OT0702-6b	<i>n</i> =2	61.27	0.31	19.01	2.89	0.13	0.58	2.05	5.5	7.81	0	0.47	100	13.31	1.42	
		<b>sd</b>	<b>0.23</b>	<b>0.24</b>	<b>0.21</b>	<b>0.18</b>	<b>0.16</b>	<b>0.16</b>	<b>0.36</b>	<b>0.59</b>	<b>0</b>	<b>0</b>	<b>0.14</b>				
	OT0520-3a	<i>n</i> =3	61.16	0.43	18.96	3.03	0.23	0.35	1.72	6.17	7.08	0.05	0.81	100	13.25	1.15	
		<b>sd</b>	<b>0.15</b>	<b>0.02</b>	<b>0.07</b>	<b>0.09</b>	<b>0.01</b>	<b>0</b>	<b>0.05</b>	<b>0.13</b>	<b>0.35</b>	<b>0.01</b>	<b>0.01</b>				
	OT0520-3b	<i>n</i> =6	61.94	0.43	19.09	3.03	0.2	0.46	1.97	4.79	7.39	0.07	0.65	100	12.18	1.54	
		<b>sd</b>	<b>0.38</b>	<b>0.04</b>	<b>0.2</b>	<b>0.07</b>	<b>0.05</b>	<b>0.13</b>	<b>0.25</b>	<b>0.43</b>	<b>0.56</b>	<b>0.04</b>	<b>0.19</b>				
	PT0704-3a	<i>n</i> =25	60.77	0.44	19.13	2.93	0.26	0.4	1.64	6.38	7.33	0	0.71	100	13.72	1.15	
		<b>sd</b>	<b>0.24</b>	<b>0.08</b>	<b>0.13</b>	<b>0.08</b>	<b>0.07</b>	<b>0.07</b>	<b>0.09</b>	<b>0.19</b>	<b>0.19</b>	<b>0</b>	<b>0.05</b>				
PT0704-3b	<i>n</i> =10	61.34	0.43	19.45	3.06	0.17	0.59	1.85	5.13	7.47	0	0.5	99.99	12.6	1.46		
	<b>sd</b>	<b>0.9</b>	<b>0.08</b>	<b>0.43</b>	<b>0.29</b>	<b>0.08</b>	<b>0.25</b>	<b>0.48</b>	<b>0.57</b>	<b>0.71</b>	<b>0</b>	<b>0.14</b>					
PT0704-3c	<i>n</i> =4	61.22	0.41	19.08	3.27	0.11	0.79	2.2	3.26	9.31	0	0.36	100	12.57	2.86		
	<b>sd</b>	<b>1.02</b>	<b>0.08</b>	<b>0.17</b>	<b>0.29</b>	<b>0.13</b>	<b>0.17</b>	<b>0.36</b>	<b>0.46</b>	<b>0.93</b>	<b>0</b>	<b>0.08</b>					
Prox.-PDCs	<i>n</i> =10	61.9	0.4	18.35	3.09	0.21	0.47	2.03	5.15	7.74	0.09	0.59	100.02	12.89	1.5		
	<b>sd</b>	<b>0.85</b>	<b>0.04</b>	<b>0.32</b>	<b>0.27</b>	<b>0.08</b>	<b>0.16</b>	<b>0.34</b>	<b>0.58</b>	<b>0.51</b>	<b>0.06</b>	<b>0.2</b>					
Prox.-Fall	<i>n</i> =10	61.33	0.42	19.16	2.89	0.24	0.35	1.75	5.96	7	0.05	0.84	99.99	12.96	1.17		
	<b>sd</b>	<b>0.5</b>	<b>0.02</b>	<b>0.22</b>	<b>0.12</b>	<b>0.03</b>	<b>0.02</b>	<b>0.06</b>	<b>0.33</b>	<b>0.19</b>	<b>0.02</b>	<b>0.08</b>					
LGM/TM18	<i>n</i> =42	61.79	0.42	19.11	2.93	0.24	0.35	1.73	5.66	6.93	0.05	0.78	99.99	12.59	1.22		
	<b>sd</b>	<b>0.46</b>	<b>0.02</b>	<b>0.21</b>	<b>0.09</b>	<b>0.02</b>	<b>0.02</b>	<b>0.07</b>	<b>0.71</b>	<b>0.29</b>	<b>0.03</b>	<b>0.05</b>					
SA3-a/C-20	OT0701-6	<i>n</i> =15	59.58	0.44	20.04	2.76	0.13	0.51	2.31	5.24	8.8	0	0.18	100	14.04	1.68	
		<b>sd</b>	<b>0.23</b>	<b>0.07</b>	<b>0.12</b>	<b>0.14</b>	<b>0.09</b>	<b>0.09</b>	<b>0.09</b>	<b>0.16</b>	<b>0.17</b>	<b>0</b>	<b>0.04</b>				
	Proximal	<i>n</i> =10	59.71	0.51	19.41	2.97	0.25	0.5	1.98	5.26	8.8	-	0.61	100	14.06	1.67	
		<b>sd</b>	<b>0.27</b>	<b>0.08</b>	<b>0.21</b>	<b>0.13</b>	<b>0.11</b>	<b>0.08</b>	<b>0.1</b>	<b>0.17</b>	<b>0.19</b>	<b>-</b>	<b>0.03</b>				
C-20	n.a.	59.45	0.49	19.96	3.07	-	0.32	2.73	4.4	9.51	-	-	99.93	13.91	2.16		

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Table 4. Continued.

			SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	FeO <sub>tot</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	P <sub>2</sub> O <sub>5</sub>	ClO	Total	Total alkali	alkali ratio
TAU1-b/X-5	OT0702-8	<i>n</i> =12	57.6	0.55	19.49	4.48	0.16	1.22	3.77	4.18	8.09	0.01	0.46	100	12.26	1.94
		<b>sd</b>	<b>0.48</b>	<b>0.08</b>	<b>0.16</b>	<b>0.34</b>	<b>0.07</b>	<b>0.16</b>	<b>0.33</b>	<b>0.19</b>	<b>0.45</b>	<b>0.03</b>	<b>0.05</b>			
	OT0701-7a	<i>n</i> =8	50.86	1.15	18.55	8.33	0.22	3.19	8.47	3.27	5.4	0.28	0.29	99.99	8.66	1.65
		<b>sd</b>	<b>1.79</b>	<b>0.16</b>	<b>0.33</b>	<b>1.26</b>	<b>0.06</b>	<b>0.59</b>	<b>1.15</b>	<b>0.45</b>	<b>0.78</b>	<b>0.14</b>	<b>0.04</b>			
	OT0701-7b	<i>n</i> =8	58.3	0.63	19.75	3.62	0.15	0.82	3.41	4.54	8.36	0.01	0.4	99.99	12.9	1.84
		<b>sd</b>	<b>1.48</b>	<b>0.12</b>	<b>0.2</b>	<b>0.89</b>	<b>0.08</b>	<b>0.43</b>	<b>1.07</b>	<b>0.31</b>	<b>0.62</b>	<b>0.03</b>	<b>0.05</b>			
	LGM/TM24a-1	<i>n</i> =32	57.23	0.51	19	4.46	0.14	1.2	4.21	4.23	8.23	0.26	0.52	99.99	12.46	1.95
		<b>sd</b>	<b>0.59</b>	<b>0.03</b>	<b>0.19</b>	<b>0.2</b>	<b>0.02</b>	<b>0.12</b>	<b>0.3</b>	<b>0.1</b>	<b>0.22</b>	<b>0.04</b>	<b>0.04</b>			
LGM/TM24a-2	<i>n</i> =6	59.48	0.41	19.14	3.6	0.14	0.73	2.94	4.26	8.62	0.13	0.55	100	12.88	2.02	
	<b>sd</b>	<b>0.45</b>	<b>0.03</b>	<b>0.16</b>	<b>0.21</b>	<b>0.03</b>	<b>0.1</b>	<b>0.26</b>	<b>0.18</b>	<b>0.22</b>	<b>0.04</b>	<b>0.05</b>				
LGM/TM24b-1	<i>n</i> =7	61.25	0.38	18.86	3.22	0.14	0.6	2.51	4.1	8.32	0.1	0.51	99.99	12.42	2.03	
	<b>sd</b>	<b>0.8</b>	<b>0.03</b>	<b>0.29</b>	<b>0.24</b>	<b>0.03</b>	<b>0.08</b>	<b>0.19</b>	<b>0.35</b>	<b>0.32</b>	<b>0.03</b>	<b>0.08</b>				
LGM/TM24b-2	<i>n</i> =10	58.39	0.53	18.77	4.6	0.15	1.29	3.9	3.87	7.82	0.26	0.42	100	11.69	2.02	
	<b>sd</b>	<b>0.86</b>	<b>0.04</b>	<b>0.37</b>	<b>0.34</b>	<b>0.02</b>	<b>0.23</b>	<b>0.47</b>	<b>0.16</b>	<b>0.42</b>	<b>0.06</b>	<b>0.06</b>				
X-6	OT0702-9	<i>n</i> =15	61.15	0.46	18.82	3.09	0.29	0.39	1.68	6.39	7.03	0	0.71	100	13.42	1.12
		<b>sd</b>	<b>0.31</b>	<b>0.07</b>	<b>0.12</b>	<b>0.14</b>	<b>0.09</b>	<b>0.11</b>	<b>0.08</b>	<b>0.64</b>	<b>0.52</b>	0	<b>0.15</b>			
	LGM/TM27-a	<i>n</i> 06	61.56	0.45	18.52	2.78	0.2	0.45	1.89	5.88	7.57	0.08	0.63	100.01	13.45	1.29
<b>sd</b>		<b>0.18</b>	<b>0.02</b>	<b>0.07</b>	<b>0.08</b>	<b>0.02</b>	<b>0.05</b>	<b>0.08</b>	<b>0.19</b>	<b>0.18</b>	<b>0.04</b>	<b>0.07</b>				
LGM/TM27-b	<i>n</i> =10	60.79	0.48	18.55	2.98	0.31	0.3	1.73	7.21	6.54	0.04	0.89	99.82	13.75	0.91	
	<b>sd</b>	<b>0.33</b>	<b>0.03</b>	<b>0.1</b>	<b>0.11</b>	<b>0.03</b>	<b>0.01</b>	<b>0.06</b>	<b>0.28</b>	<b>0.16</b>	<b>0.03</b>	<b>0.08</b>				
P-11	OT0702-10a	<i>n</i> =3	66.18	0.43	16.23	4.14	0.23	0.2	0.84	5.85	5.79	0	0.13	100.01	11.64	0.99
		<b>sd</b>	<b>0.89</b>	<b>0.16</b>	<b>3.11</b>	<b>3.16</b>	<b>0.2</b>	<b>0.21</b>	<b>0.12</b>	<b>0.79</b>	<b>1.39</b>	0	<b>0.11</b>			
	OT0702-10b	<i>n</i> =6	72.52	0.41	9.34	6.7	0.31	0.12	0.34	5.06	4.5	0	0.71	100	9.55	0.91
		<b>sd</b>	<b>0.79</b>	<b>0.09</b>	<b>1.23</b>	<b>0.4</b>	<b>0.06</b>	<b>0.06</b>	<b>0.07</b>	<b>0.71</b>	<b>0.31</b>	0	<b>0.08</b>			
	KET82-22	<i>n</i> a.	64.67	0.83	15.5	5.88	0	0.28	1.45	6.41	4.98	0	0	100	11.39	0.78
		<b>sd</b>	<b>0.68</b>	<b>0.1</b>	<b>0.71</b>	<b>0.19</b>	0	<b>0.14</b>	<b>0.2</b>	<b>0.27</b>	<b>0.19</b>	0	<b>0</b>			
	Lesvos/ML5-a	<i>n</i> =10	73.73	0.43	8.31	7.14	0.32	0.08	0.31	4.3	4.44	0.01	0.84	99.91	8.74	1.03
		<b>sd</b>	<b>0.24</b>	<b>0.01</b>	<b>0.1</b>	<b>0.16</b>	<b>0.02</b>	<b>0.01</b>	<b>0.01</b>	<b>0.25</b>	<b>0.09</b>	<b>0.01</b>	<b>0.03</b>			
Lesvos/ML5-b	<i>n</i> =15	66.23	0.75	14.32	6.06	0.3	0.27	1.16	5.59	4.99	0.1	0.17	99.94	10.58	0.89	
	<b>sd</b>	<b>1.04</b>	<b>0.06</b>	<b>1.1</b>	<b>0.48</b>	<b>0.02</b>	<b>0.1</b>	<b>0.27</b>	<b>0.28</b>	<b>0.14</b>	<b>0.04</b>	<b>0.06</b>				

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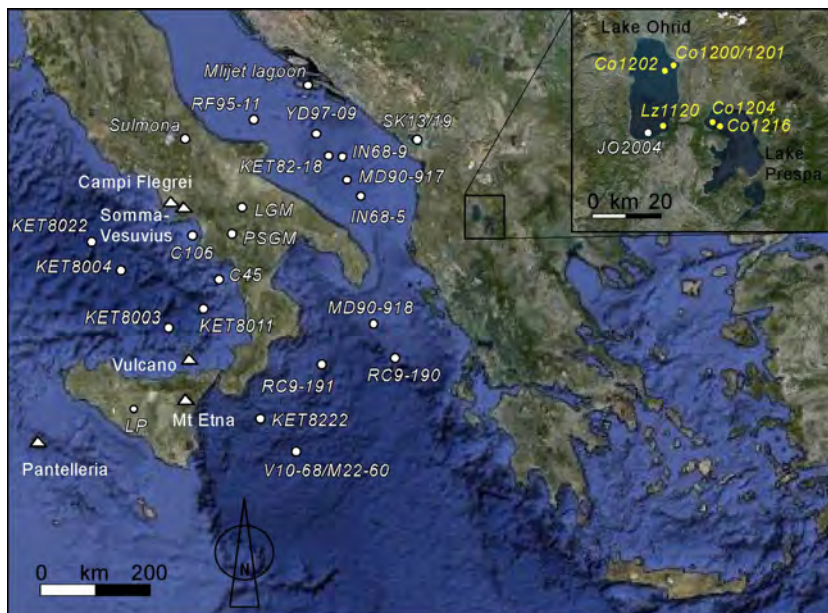
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**Fig. 1.** Location map of the study area. A close view of lakes Ohrid and Prespa is shown in the framework in the upper right angle, along with the locations of the studied cores. Triangles indicate Italian volcanoes of interest for this study. White dots indicate the cores used for tephra correlations: LP, Lago di Pergusa (Sadori and Narcisi, 2001); JO2004, Albanian side of Lake Ohrid (Caron et al., 2010a; Lézine et al., 2010); LGM, Lago Grande di Monticchio (Wulf et al., 2004), Sulmona (Giaccio et al., 2009), YD97-09, RF95-11 (Lowe et al., 2007); IN68-5, IN68-9 (Calanchi and Dinelli, 2008); KET82-18, KET80-03, KET80-04, KET80-011, KET80-22 (Paterne et al., 1988); KET82-22 (Paterne et al., 2008); MD90-917 (Siani et al., 2004); MD90-918 (Caron et al., 2010b; Zanchetta et al., 2010); C45, C146 (Buccheri et al., 2002); V10-68, M22-60, RC9-190, RC9-191 (Keller et al., 1978); PSGM, Pantano San Gregorio Magno (Munno and Petrosino, 2006); Mlijet Lagoon (Jahns and van den Boogard, 1998); SK13/19, Shkodra lake (Sulpizio et al., 2010).

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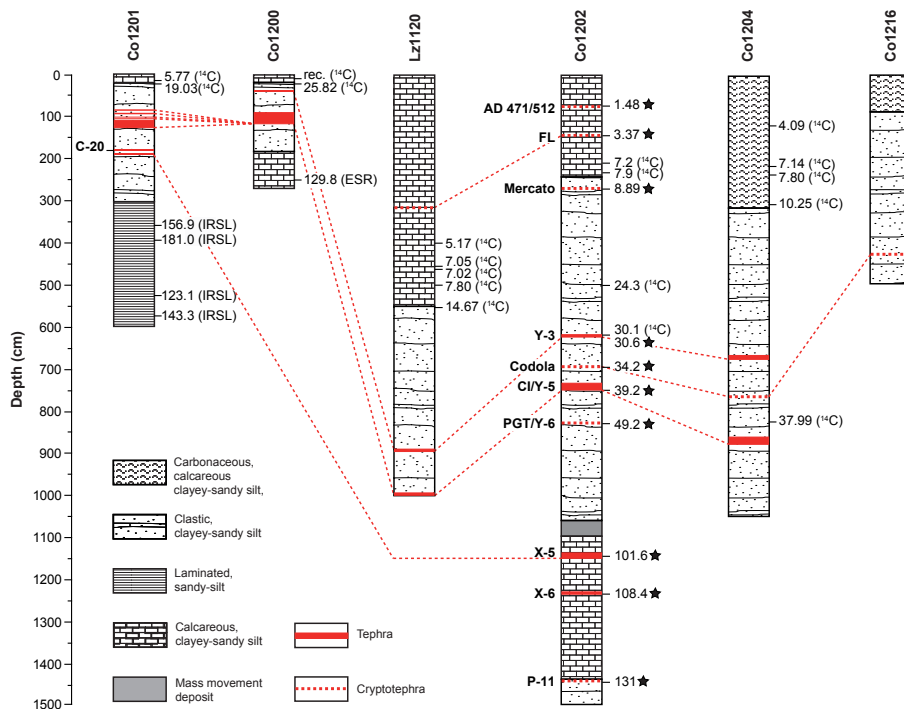
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**Fig. 2.** Summary of stratigraphy and lithology of the studied cores. The thicknesses and sediment depths of tephtras and cryptotephtras are indicated by red horizons. Dashed lines between the cores indicate the correlation of individual tephtras and cryptotephtras. Age control for the different sediment successions apart from tephrochronology is given by radiocarbon (<sup>14</sup>C), infrared stimulated luminescence (IRSL), and electron spin resonance (ESR) dating.

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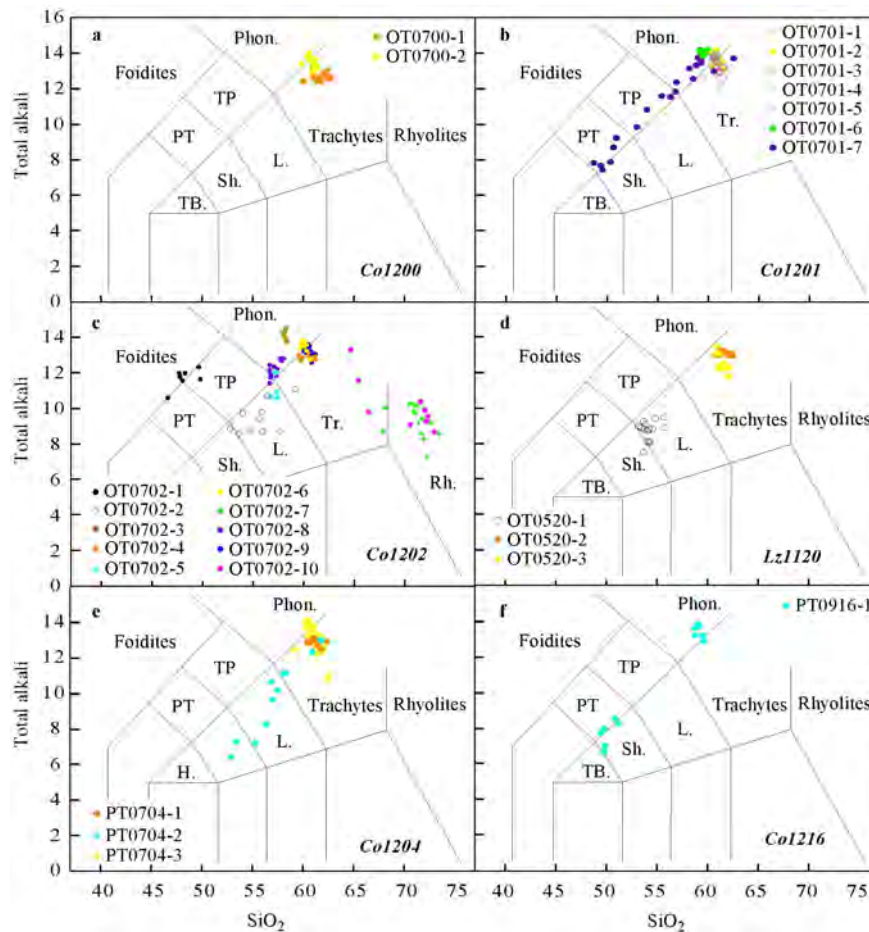
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**Fig. 3.** Total alkali vs. silica diagrams for the tephra layers recognised in the six cores studied or reviewed in this paper.

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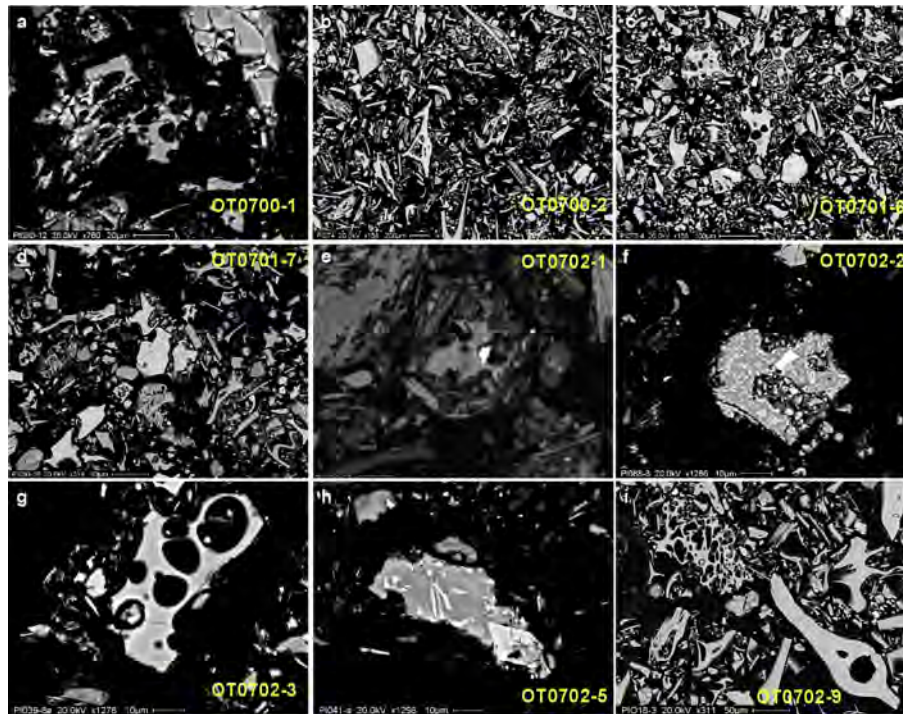
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**Fig. 4.** SEM photographs of volcanic fragments from selected tephra layers and cryptotephra. **(a)** Y-3 tephra layer in core Co1200; **(b)** CI/Y-5 tephra layer from core Co1200; **(c)** C-20 cryptotephra in core Co1201; **(d)** X-5 tephra layer in core Co1201; **(e)** AD 472 cryptotephra in core Co1202; **(f)** FL cryptotephra in Co 1202 core; **(g)** Mercato cryptotephra in Co1202 core; **(h)** Codola cryptotephra in core Co1202; **(i)** X-6 tephra layer core Co1202.

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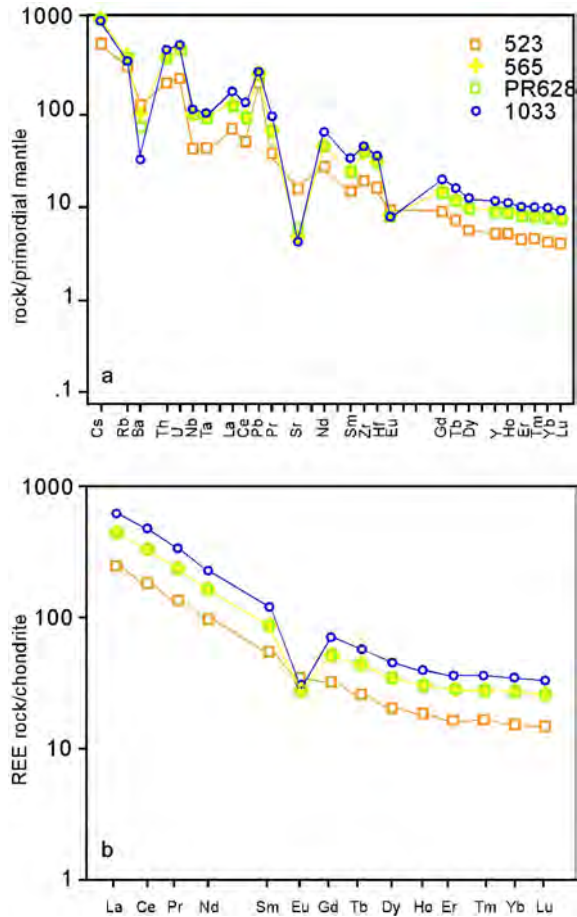
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**Fig. 5.** Primitive-mantle and chondrite-normalised rare earth element (REE) diagrams for the selected samples. Samples 523, 535 and 1033 are from core Co1202, the sample PR628 is from core Co1204.

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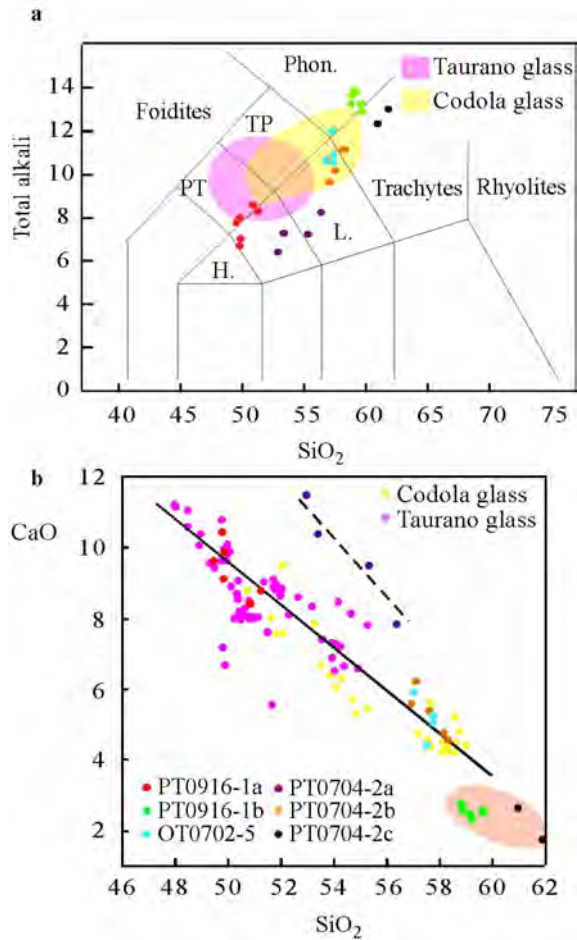
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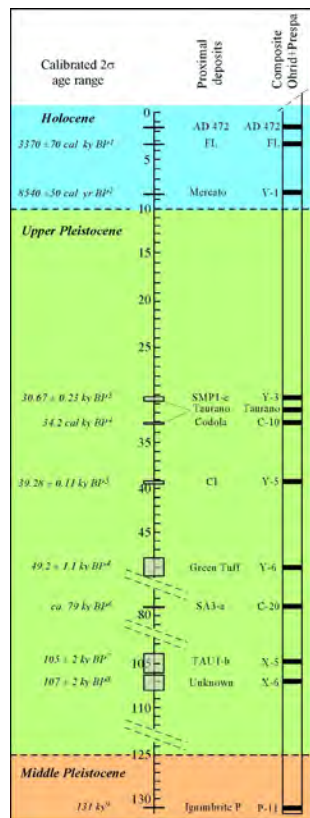
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**Fig. 6.** (a) Total alkali vs. silica diagram for the different groups of OT0702-5 (pale blue dots), PT09-16 (red and green dots), PT0704-2 (purple, orange and black dots); (b) CaO vs. SiO<sub>2</sub> diagram for the same groups of Fig. 5a.



**Fig. 7.** Updated tephrochronology for the last 131 ky of lakes Ohrid and Prespa. Grey boxes indicate the age uncertainties for calibrated  $^{14}\text{C}$  data and for  $^{39}\text{Ar}/^{40}\text{Ar}$  ages. <sup>1</sup> Coltelli et al. (2000); <sup>2</sup> Zanchetta et al. (2010); <sup>3</sup> Di Vito et al. (2008); <sup>4</sup> Vogel et al. (2010); <sup>5</sup> De Vivo et al. (2001); <sup>6</sup> Paterne et al. (1988); <sup>7</sup> Kraml (1997); <sup>8</sup> Allen et al. (1999); <sup>9</sup> Paterne et al. (2008). The three colours indicate the Middle Pleistocene, Upper Pleistocene and Holocene.

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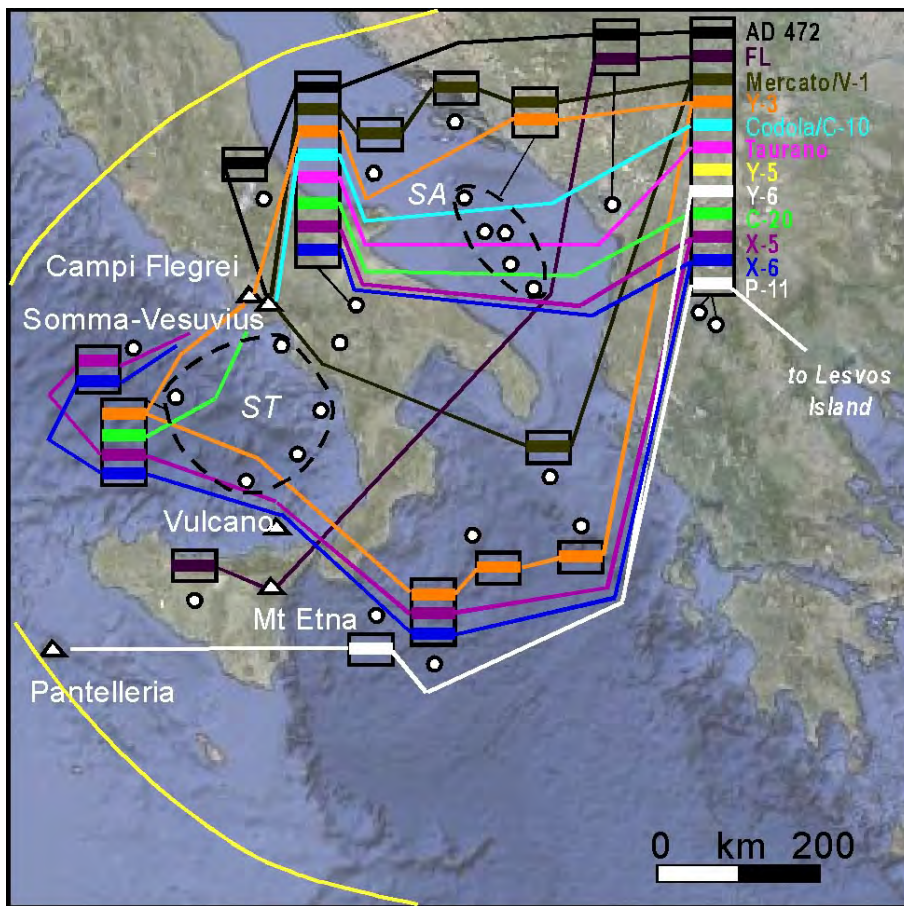
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**Fig. 8.** Link of the Ohrid and Prespa tephrostratigraphy to the other archives of the Central Mediterranean area. SA=South Adriatic Sea; ST=South Tyrrhenian Sea.

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