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# Soil CO<sub>2</sub> and N<sub>2</sub>O emissions and microbial abundances altered by temperature rise and nitrogen addition in active-layer soils of permafrost peatland

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Changes in soil CO<sub>2</sub> and N<sub>2</sub>O emissions due to climate change and nitrogen input will result in increased levels of atmospheric CO<sub>2</sub> and N<sub>2</sub>O, thereby feeding back into Earth's climate. Understanding the responses of soil carbon and nitrogen emissions mediated by microbe from permafrost peatland to temperature rising is important for modeling the regional carbon and nitrogen balance. This study conducted a laboratory incubation experiment at 15 and 20°C to observe the impact of increasing temperature on soil CO<sub>2</sub> and N<sub>2</sub>O emissions and soil microbial abundances in permafrost peatland. An NH<sub>4</sub>NO<sub>3</sub> solution was added to soil at a concentration of 50mgNkg<sup>-1</sup> to investigate the effect of nitrogen addition. The results indicated that elevated temperature, available nitrogen, and their combined effects significantly increased CO<sub>2</sub> and N<sub>2</sub>O emissions in permafrost peatland. However, the temperature sensitivities of soil CO<sub>2</sub> and N<sub>2</sub>O emissions were not affected by nitrogen addition. Warming significantly increased the abundances of methanogens, methanotrophs, and *nirK*-type denitrifiers, and the contents of soil dissolved organic carbon (DOC) and ammonia nitrogen, whereas *nirS*-type denitrifiers, β-1,4-glucosidase (βG), cellobiohydrolase (CBH), and acid phosphatase (AP) activities significantly decreased. Nitrogen addition significantly increased soil *nirS*-type denitrifiers abundances, β-1,4-N-acetylglucosaminidase (NAG) activities, and ammonia nitrogen and nitrate nitrogen contents, but significantly reduced bacterial, methanogen abundances, CBH, and AP activities. A rising temperature and nitrogen addition had synergistic effects on soil fungal and methanotroph abundances, NAG activities, and DOC and DON contents. Soil CO<sub>2</sub> emissions showed a significantly positive correlation with soil fungal abundances, NAG activities, and ammonia nitrogen and nitrate nitrogen contents. Soil N<sub>2</sub>O emissions showed positive correlations with soil fungal, methanotroph, and *nirK*-type denitrifiers abundances, and DOC, ammonia nitrogen, and nitrate contents. These results demonstrate the importance of soil microbes, labile carbon, and nitrogen for regulating soil carbon and nitrogen emissions. The results

of this study can assist simulating the effects of global climate change on carbon and nitrogen cycling in permafrost peatlands.

#### KEYWORDS

climate warming, nitrogen availability, soil microbial abundance, enzyme activity, boreal peatland

## Introduction

Soil carbon dioxide (CO<sub>2</sub>) emissions represent the second largest carbon (C) flux in terrestrial ecosystems, accounting for 70–90% of total ecosystem respiration (Schlesinger and Andrews, 2000; Cascio et al., 2017). Losses of soil C to the atmosphere through soil heterotrophic respiration play an important role in regulating atmospheric CO<sub>2</sub>. These losses are predicted to increase due to climate change, resulting in a positive C-climate feedback loop (Yuan et al., 2019; Dacal et al., 2022). The availability of nitrogen (N) changes the source-sink dynamics of ecosystem C by changing the soil CO<sub>2</sub> flux (Cascio et al., 2017). Soils also act as an important source-sink for nitrous oxide (N<sub>2</sub>O; Wu et al., 2013, 2015). Climate warming and the input of N could change mineralization of soil N and N<sub>2</sub>O emissions (Ma et al., 2011). The increases in N<sub>2</sub>O emissions can cause changes in global warming potential, thus affecting the C sinks and CO<sub>2</sub> emissions (Muhammad et al., 2022). However, little is known about how increases in temperature and N inputs interact to regulate soil emissions of CO<sub>2</sub> and N<sub>2</sub>O and their temperature sensitivity. An increased comprehension of the microbial mechanisms under warming and N addition impact emissions of CO<sub>2</sub> and N<sub>2</sub>O is vital for accurately simulating the consequences of a changing global climate on the C and N balance.

Low temperatures and nutrient concentration limited soil microbial activities and soil organic matter (SOM) decomposition (Koyama et al., 2014). An increase in temperature results in enhanced microbial growth and in the activation of the functional genes involved in C and N cycling (Xue et al., 2016; Wang et al., 2019). These result in increased soil C decomposition and respiration (Han et al., 2013). However, a previous study noted a reduction in N<sub>2</sub>O production with increasing temperature, especially due to denitrification (Duan et al., 2019), whereas the abundances of *amoA*, *nifH*, and *nirK* increased (Jung et al., 2011; Han et al., 2013). Warming could increase N limitation of microorganisms, which, in turn, could limit the impact of increased temperature on SOM mineralization. Previous studies found that N addition increased the abundances of C decomposition and N cycling genes (Jung et al., 2011; Wang et al., 2019), leading to a stronger positive correlation between soil available N and microbial properties exposed to elevated temperature (Huang et al., 2022). Greater insight into the impacts of warming and the addition of N on soil microorganisms can

assist in improving understanding of the reactions of soil C and N emissions to a global changing climate.

Soil enzymes catalyze breakdown of high molecular weight compounds, and play important functions in SOM degradation (Yao et al., 2015), measuring their activities can provide useful indicators of soil emissions of CO<sub>2</sub> and N<sub>2</sub>O (Chen et al., 2017). Soil enzyme activities can be used to investigate microbial nutrient cycling due to their connections with active microbial biomass, including microbial responses to environmental changes, transformation rates, and the location of the most active biomass (Wang et al., 2015). Warming can result in changes in enzyme activities, leading to functional changes in soil ecosystem processes (Xu et al., 2015). An improved understanding of decomposition and mechanisms of microbial enzyme production can assist in constraining long-term responses to warming (Sihl et al., 2016). Moreover, enzyme activities were applied as indicators of the impacts of N input within many recent experiments since they reflect the metabolic needs of soil microbial communities relative to available nutrients (Ochoa-Hueso et al., 2013). Nitrogen addition significantly stimulated activities of N- and phosphorus-acquiring hydrolytic enzymes and depressed the activities of oxidative enzymes (Tu et al., 2014). Maslov and Maslova (2021) investigated the effect of increased N availability on changes in soil enzyme activities to better understand the internal mechanisms of soil C and N cycling processes. Improved comprehension of soil enzymes and their regulatory mechanisms is needed to enhance comprehension of the impacts of temperature and N availability on soil CO<sub>2</sub> and N<sub>2</sub>O emissions.

Peatlands represent an important C pool on Earth, storing 1,055 Gt of soil C, even though they only cover 3% of the land surface of the Earth (Nichols and Peteet, 2019). In particular, permafrost peatlands experience increased storage and emissions of C, and can act as key contributors to global warming. Permafrost thaw in northern peatlands results in alterations to ground thermal conditions, moisture, and chemistry, which, in turn, regulate microbial activities responsible for generating greenhouse gases (GHGs) from decomposing organic matter (Kirkwood et al., 2021). Newly thawed permafrost in Western Canada is predicted to release 0.2 to 25% of stored C by 2,100 (Jin and Ma, 2021). An increase in annual temperature by 1°C was predicted to increase respiration by up to 60% in an experiment conducted in Arctic blanket peatland (Dorrepaal et al., 2009). Moreover, increases in N input affected N<sub>2</sub>O emissions in northern peatlands due to increased N availability and/or changing vegetation composition

(Le et al., 2020). Nitrogen addition could mitigate the positive effect of warming on methane fluxes in a coastal bog (Gong et al., 2021). However, the synergistic environmental parameters regulating GHGs emissions in northern permafrost peatlands remain largely unknown (AminiTabrizi et al., 2020). Clarifying the synergistic effects of both climate warming and a rising nitrogen availability on permafrost emissions of CO<sub>2</sub> and N<sub>2</sub>O can provide a reference for future studies on potential responses of C and N sequestration of high latitude peatlands to climate change.

Northeastern China contains the second largest expanse of permafrost in China, primarily known as Xing'an-Baikal permafrost. This permafrost area lies on the southeastern edges of the Eurasian cryolithozone and is thermally unstable and sensitive to external changes (Wei et al., 2011). By the 2010s, the area of Xing'an-Baikal permafrost in Northeast China had declined by 40.6% compared with that in the 1960s (Li et al., 2021). The present study aimed to understand the synergistic effects of both climate warming and rising N availability on soil emissions of CO<sub>2</sub> and N<sub>2</sub>O and its regulation mechanism in permafrost peatlands. An incubation experiment with temperature increase of 5°C and nitrogen addition of 50 mg N kg<sup>-1</sup> was conducted in the Great Xing'an mountain peatland, Northeast China. The objectives of this research were to explore the response of CO<sub>2</sub> and N<sub>2</sub>O emissions from permafrost peatland soil to warming and nitrogen addition, and clarify their driving mechanisms, which can help improve future predictions of responses of soil C and N cycling to climate warming.

## Materials and methods

### Site description and soil sampling

The study site of the present study is a typical permafrost peatland nearby the Tuqiang Forestry Bureau, Great Xing'an Mountain (52°44'N, 122°39'E), Heilongjiang Province, China. Average yearly temperature and average yearly precipitation are -3.9°C and 452 mm, respectively. The dominant species of plants are *Vaccinium uliginosum* L., Moench, *Sphagnum* spp., *Ledum palustre* L., *Eriophorum vaginatum* L., and *Chamaedaphne calyculata* L. The soil type of the study area according to the United States Department of Agriculture (USDA) classification system is Glacic Histoturbels (Soil Survey Staff, 2010). A soil sample of the active layer (0–20 cm) was obtained using a hand auger soil core sampler, which was filtered through a 2-mm sieve. The total C (TC) and total N (TN) of the soil sample before incubation experiments were 408.74 and 15.34 g kg<sup>-1</sup>, respectively, whereas soil moisture and pH were 77.18% and 5.49, respectively.

### Laboratory incubation

Fresh soil samples (15 g according to completely dry soil) were placed in 500-ml glass flasks and preincubated at 15°C for 7 days.

NH<sub>4</sub>NO<sub>3</sub> solution (2 ml) was uniformly added to soil at a concentration of 50 mg N kg<sup>-1</sup>, with four replicates prepared. Deionized water (2 ml) was added to the control treatment. The flask lids were sealed with rubber septa to allow the analysis of rates of emissions of CO<sub>2</sub> and N<sub>2</sub>O at 15 and 20°C (maximum monthly mean temperature in July of 18.4°C). These soils were incubated continuously for 18 days. Trapped air in the jars was removed for CO<sub>2</sub> and N<sub>2</sub>O determination at intervals of 2 h, 1, 2, 3, 5, 7, 9, 12, 15, and 18 days. Headspace gas in the jars was extracted using a 50-ml syringe with a three-way valve. The concentrations of CO<sub>2</sub> and N<sub>2</sub>O were measured utilizing a gas chromatograph (Agilent 7890B, United States). Deionized water corresponding to the reduction in weight after each collection of gas was added. Soil samples were collected to determine soil microbial abundances, enzyme activities, and labile C and N contents at the end of incubation.

### Soil microbial abundances analysis

Soil DNA was extracted from a 300-mg subsample using a FastDNA spin Kit (MPbio, Santa Ana, CA, United States) in accordance with the manufacturer's instructions. Bacterial 16S rRNA, fungal IST, and functional genes encoding *mcrA*, *pmoA*, *nirS*, and *nirK* were quantitatively evaluated *via* qPCR using an ABI StepOne instrument (Applied Biosystems, San Francisco, CA, United States). Supplementary Table S1 lists the primers and amplification details used in the present study. The PCR mixture contained 10 ng soil DNA, 0.4 μl primers (10 μM), and 12.5-μl of SYBR Buffer (TaKaRa, Beijing, China) in a final volume of 25 μl. qPCR standard curves were created by purifying amplicon products of functional and phylogenetic markers using a cyclic purification kit (Omega Bio-Tek, United States), ligated to the pMD18-T (TaKaRa) vector, and transforming into *Escherichia coli*. A plasmid mini kit (Omega Bio-Tek, United States) was utilized to remove the plasmids, with a standard local alignment searching tool used to identify specificity of plasmids. Standard curves were produced by plasmid serial dilution (Song et al., 2021).

### Soil enzyme activities measurement

The potential activities of acid phosphatase (AP), β-1,4-glucosidase (βG), cellobiohydrolase (CBH), and NAG were measured for absorbance using a microplate spectrophotometer. Aliquots (200 μl) of slurry (1 g fresh soil sample homogenized in 125-ml 50-mM acetate buffer, pH 8) and 50-μl of substrate solution (200 μM) were placed into 96-well microplates. Every microplate had eight replicate wells per assay, as well as negative and positive controls for quench correction. The microplates were incubated in darkness at 20°C for 4 h. Excitation and emission fluorescence were identified at 365 and 450 nm, respectively using Cell Imaging Multi-Mode Reader (BioTek Cytation 5, United States).

## Soil carbon and nitrogen content measurement

Soil ammonia nitrogen ( $\text{NH}_4^+\text{-N}$ ), nitrate ( $\text{NO}_3^-\text{-N}$ ), and dissolved organic N (DON) were extracted through the addition of 2 M KCl at a 1:15 ratio, followed by 1 h of shaking at 150 rpm at a temperature of 20°C. DON concentrations of soil were calculated as the difference between total dissolved N and inorganic N. Soil dissolved organic C (DOC) contents were analyzed using a Multi N/C 2100 analyzer (Analytik Jena AG, Germany) after extracting fresh soil with a 2 M KCl solution. Soil TN contents were analyzed after digestion with sulfuric acid ( $\text{H}_2\text{SO}_4$ ) and potassium sulfate ( $\text{K}_2\text{SO}_4$ ), with cupric sulfate ( $\text{CuSO}_4$ ) used as a catalyst. The products of digestion were subsequently analyzed using an AA3 continuous flow chemical analyzer (Seal Analytical, Germany). Quantification of soil moisture was by oven drying of fresh soil at 105°C to a constant weight. The pH of soil was measured using a 1:10 soil-deionized water slurry.

## Data analyses

Statistical analyses were performed in the SPSS 24.0 package. Results are shown as the average  $\pm$  standard error. A two-way analysis of variance (ANOVA) was performed to evaluate the interactions between increasing temperature and addition of N on soil emissions of  $\text{CO}_2$  and  $\text{N}_2\text{O}$ , microbial abundances, enzyme activities, and contents of soil C and N. Linear regression analysis was conducted to explore relationships between the soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions and soil microbial abundances, enzyme activities, and soil C and N contents.

The temperature sensitivities ( $Q_{10}$ ) of soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emission rates per 10°C were calculated as follows:

$$Q_{10} = \left( \frac{K_2}{K_1} \right)^{\frac{10}{T_2 - T_1}}$$

where  $T_1$  and  $T_2$  is the incubation temperatures for 15 and 20°C, respectively.  $K_1$  and  $K_2$  is the  $\text{CO}_2$  ( $\text{mg CO}_2\text{-C kg}^{-1} \text{d}^{-1}$ ) and  $\text{N}_2\text{O}$  ( $\mu\text{g N}_2\text{O-N kg}^{-1} \text{d}^{-1}$ ) emission rates at 15 and 20°C, respectively.

## Results

### Emissions of soil $\text{CO}_2$ and $\text{N}_2\text{O}$ and their sensitivity to temperature

An increase in temperature significantly stimulated soil emissions of  $\text{CO}_2$  and  $\text{N}_2\text{O}$  in the permafrost peatlands (Figures 1A,B). Soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions in the control increased by 53.57 and 45.50% at 20°C compared to that at 15°C, respectively. The addition of N resulted in increases in  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions by 52.34 and 54.53% at 20°C compared to that at

15°C, respectively. The cumulative  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions were significantly higher under N addition than that in the control at 15°C and 20°C (Figures 2A,B). The increase in cumulative  $\text{N}_2\text{O}$  emissions after the addition of N was significantly higher than the increase in  $\text{CO}_2$ . There were significant interactions between rising temperature and addition of N on both  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions ( $p < 0.05$ ; Table 1). The sensitivities of soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions to temperature in the control were 2.37 and 2.36, respectively. The addition of N did not impact the  $Q_{10}$  values of  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions of 2.50 and 2.44, respectively (Figure 2C).

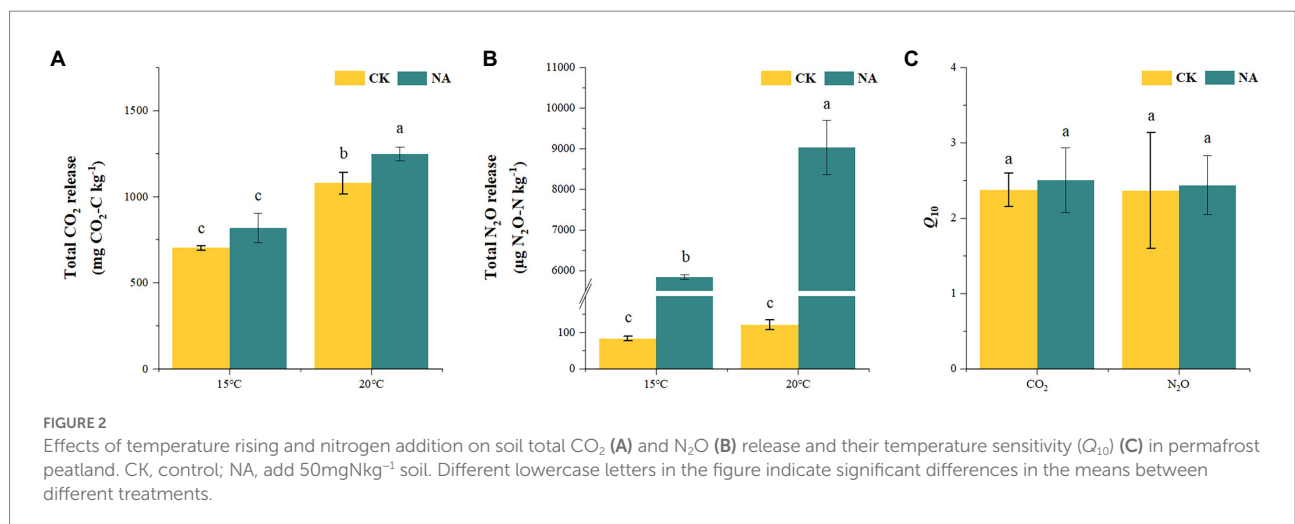
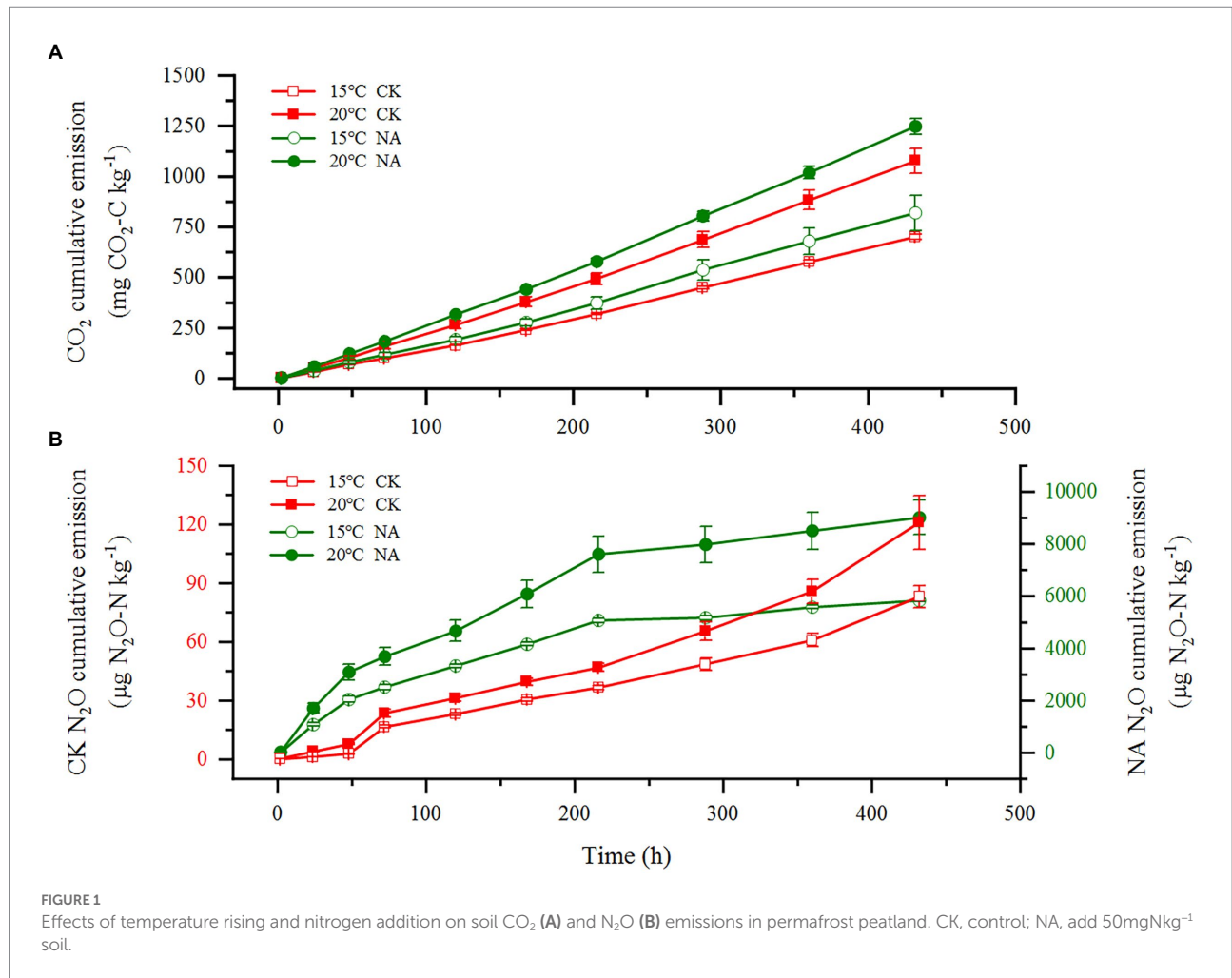
### Soil microbial abundances

Among the microbial community, bacteria were the most abundant ( $6.08\text{--}14.52 \times 10^{12}$  copies  $\text{g}^{-1}$  dry soil). At 20°C, bacterial abundances in the control and N addition treatment decreased to 36.89 and 50.54% of that at 15°C (Figure 3A), respectively, indicating the preference of bacteria for lower temperature. At 20°C, fungal abundances increased significantly by 60.73% in the N addition treatment (Figure 3B). N addition appeared to reduce the abundances of bacteria under both temperatures, whereas fungal abundances were significantly stimulated at 20°C. Increased temperature resulted in the proliferation of methanogen (*mcrA*) by 28.04 and 31.46% in the control and N addition treatments, respectively (Figure 3C). However, N addition reduced methanogen abundances by 19.30 and 17.14% at 15 and 20°C, respectively. The abundances of methanotrophs (*pmoA*) significantly increased by 28.49-, 14.31-, and 18.16-fold under a rising temperature, N addition, and both increased temperature and N addition, respectively (Figure 3D). Adding N at 15°C significantly increased the abundances of *nirK*-type denitrifiers by 21.89% (Figure 3E). An increase in temperature resulted in decreases in the abundances of the *nirS*-type denitrifiers by 25.59 and 22.75% in the control and N addition treatments, respectively (Figure 3F). The addition of N resulted in increases in the abundances of *nirS*-type denitrifiers by 19.48 and 24.04% at 15 and 20°C, respectively. The increase in temperature and N addition had an interactive impact on the abundances of fungi and methanotrophs; however, there was no synergistic effect on bacterial, methanogen, and denitrifier abundances ( $p < 0.01$ ; Table 2). There were significant relationships between the abundances of fungi and the contents of  $\text{NH}_4^+\text{-N}$ ,  $\text{NO}_3^-\text{-N}$ , as well as emissions of  $\text{CO}_2$ . This result indicated that fungi contributed to  $\text{CO}_2$  emissions and were affected by N concentrations. The significant correlations between  $\text{N}_2\text{O}$  emissions and the abundances of fungi, methanotrophs, and *nirK*-type denitrifiers indicated the significant contribution of the microbial community to  $\text{N}_2\text{O}$  emissions ( $p < 0.05$ ; Figure 4).

### Soil enzymes activities

The activities of the four soil enzymes responded significantly to a rising temperature and the addition of N (Figure 5). The





C-cycling-related activities of  $\beta$ G and CBH decreased by 22.63 and 22.46% with a rising in temperature in the control, whereas they decreased by 12.40 and 46.03% in the N addition treatment, respectively (Figures 5A,B). The rise in temperature resulted in an

increase in soil NAG activities by 11.83 and 48.57% in the control and N addition treatments, respectively (Figure 5C). Significant interactive effects were observed between the rising temperature and addition of N on soil NAG activities ( $p < 0.01$ ; Table 2). NAG

TABLE 1 Two-way ANOVA of effects of temperature rising and nitrogen addition on soil CO<sub>2</sub>, N<sub>2</sub>O release, and soil carbon and nitrogen contents.

	CO <sub>2</sub> emission rate	N <sub>2</sub> O emission rate	DOC	DON	NH <sub>4</sub> <sup>+</sup> -N	NO <sub>3</sub> <sup>-</sup> -N
Temperature rising	49.824**	23.030**	8.890*	0.190	65.101**	6.160*
Nitrogen addition	6.427*	476.718**	0.003	13.075**	324.349**	11.814**
Temperature	0.217*	21.961**	4.935*	31.132**	2.469	1.985
rising×Nitrogen addition						

DOC, dissolved organic carbon; DON, dissolved organic nitrogen. \**P*<0.05; \*\**P*<0.01.

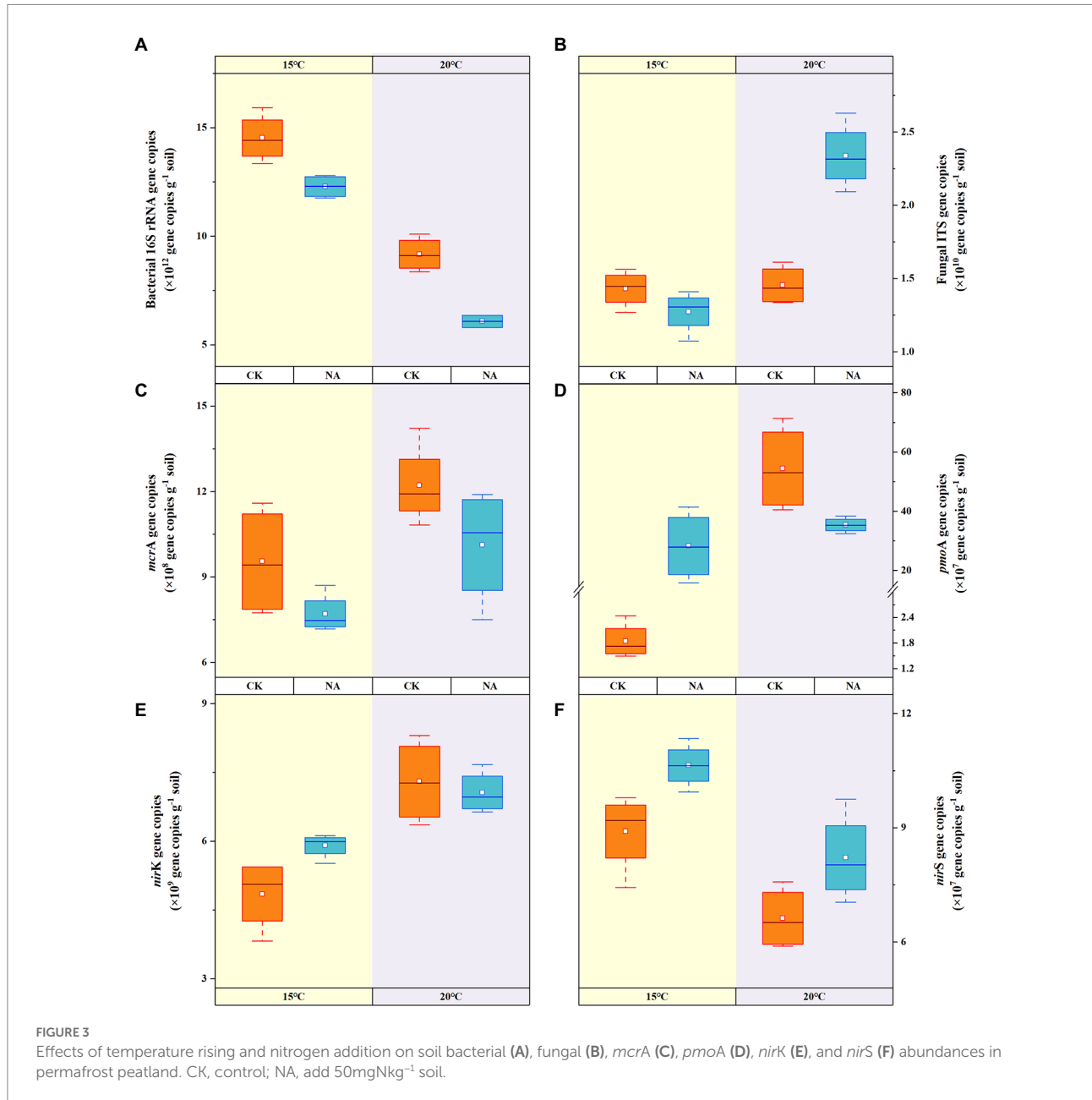


FIGURE 3 Effects of temperature rising and nitrogen addition on soil bacterial (A), fungal (B), *mcrA* (C), *pmoA* (D), *nirK* (E), and *nirS* (F) abundances in permafrost peatland. CK, control; NA, add 50mgNkg<sup>-1</sup> soil.

activities showed significant positive correlations with soil emissions of CO<sub>2</sub> and contents of NO<sub>3</sub><sup>-</sup>-N and NH<sub>4</sub><sup>+</sup>-N (*p*<0.05; Figure 4). Soil AP activities decreased with a rising temperature and the addition of N, with the highest and lowest activities of

2,089.23 and 1,730.22 nmolg<sup>-1</sup> h<sup>-1</sup> obtained at 15°C without N addition and 20°C with N addition, respectively (Figure 5D). There were no synergistic effects of the rising temperature and addition of N on soil βG, CBH, and AP activities (*p*>0.05; Table 2).

TABLE 2 Two-way ANOVA of the effects of nitrogen addition and temperature rising on soil microbial abundances and enzyme activities.

	Bacteria	Fungi	<i>mcrA</i>	<i>pmoA</i>	<i>nirK</i>	<i>nirS</i>	$\beta$ G	CBH	NAG	AP
Temperature rising	241.018**	45.530**	9.914**	39.356**	30.183**	25.749**	8.811*	23.166**	1.936	20.320**
Nitrogen addition	51.150**	20.357**	5.907*	0.597	1.564	12.909**	3.625	42.317**	67.208**	13.070**
Temperature rising $\times$ Nitrogen addition	1.308	41.716**	0.024	22.850**	3.897	0.024	1.226	0.630	19.659**	0.027

$\beta$ G,  $\beta$ -1,4-glucosidase; CBH, cellobiohydrolase; NAG,  $\beta$ -1,4-N-acetylglucosaminidase; AP, acid phosphatase. \* $P < 0.05$ ; \*\* $P < 0.01$ .

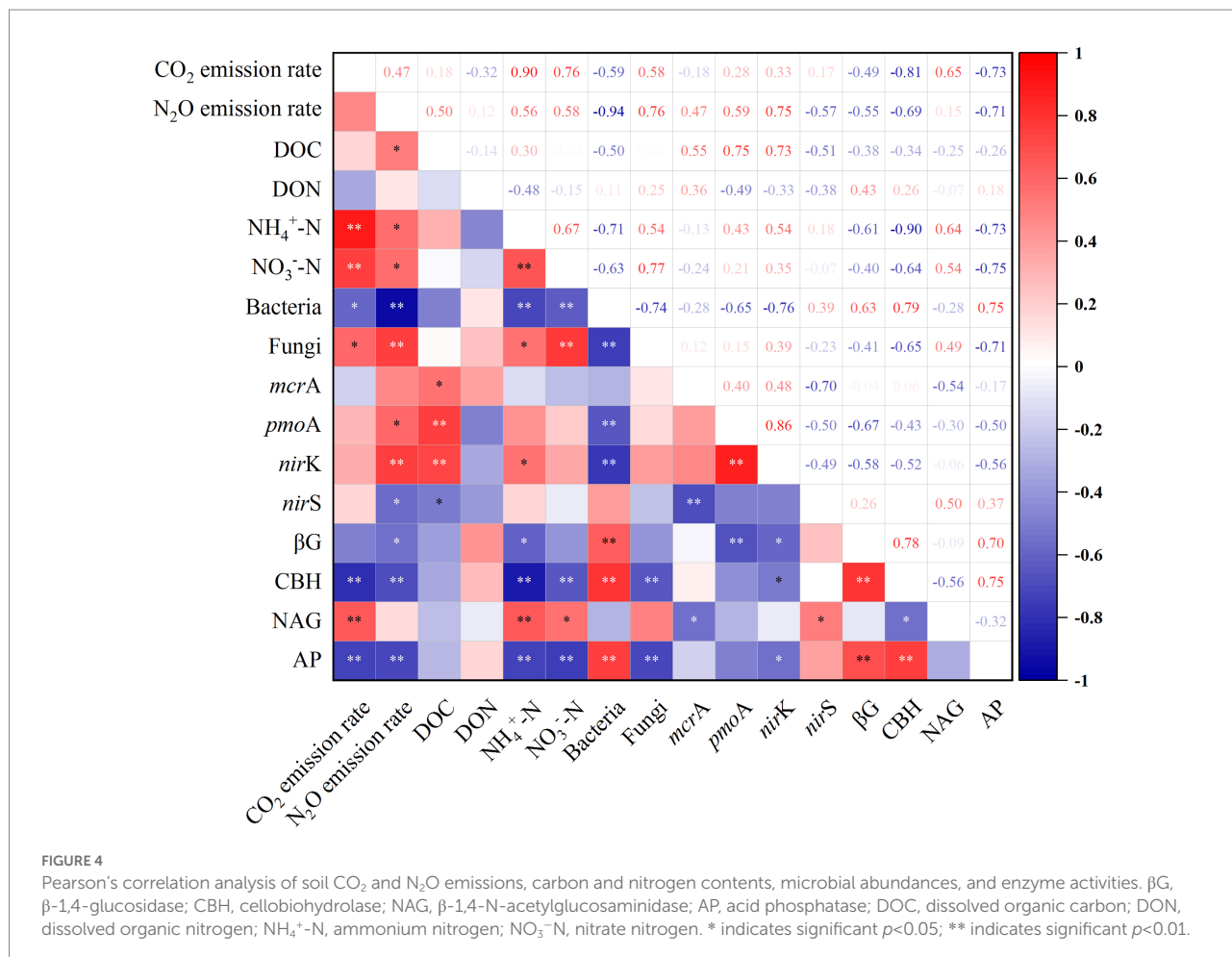


FIGURE 4

Pearson's correlation analysis of soil CO<sub>2</sub> and N<sub>2</sub>O emissions, carbon and nitrogen contents, microbial abundances, and enzyme activities.  $\beta$ G,  $\beta$ -1,4-glucosidase; CBH, cellobiohydrolase; NAG,  $\beta$ -1,4-N-acetylglucosaminidase; AP, acid phosphatase; DOC, dissolved organic carbon; DON, dissolved organic nitrogen; NH<sub>4</sub><sup>+</sup>-N, ammonium nitrogen; NO<sub>3</sub><sup>-</sup>-N, nitrate nitrogen. \* indicates significant  $p < 0.05$ ; \*\* indicates significant  $p < 0.01$ .

### Soil labile carbon and nitrogen contents

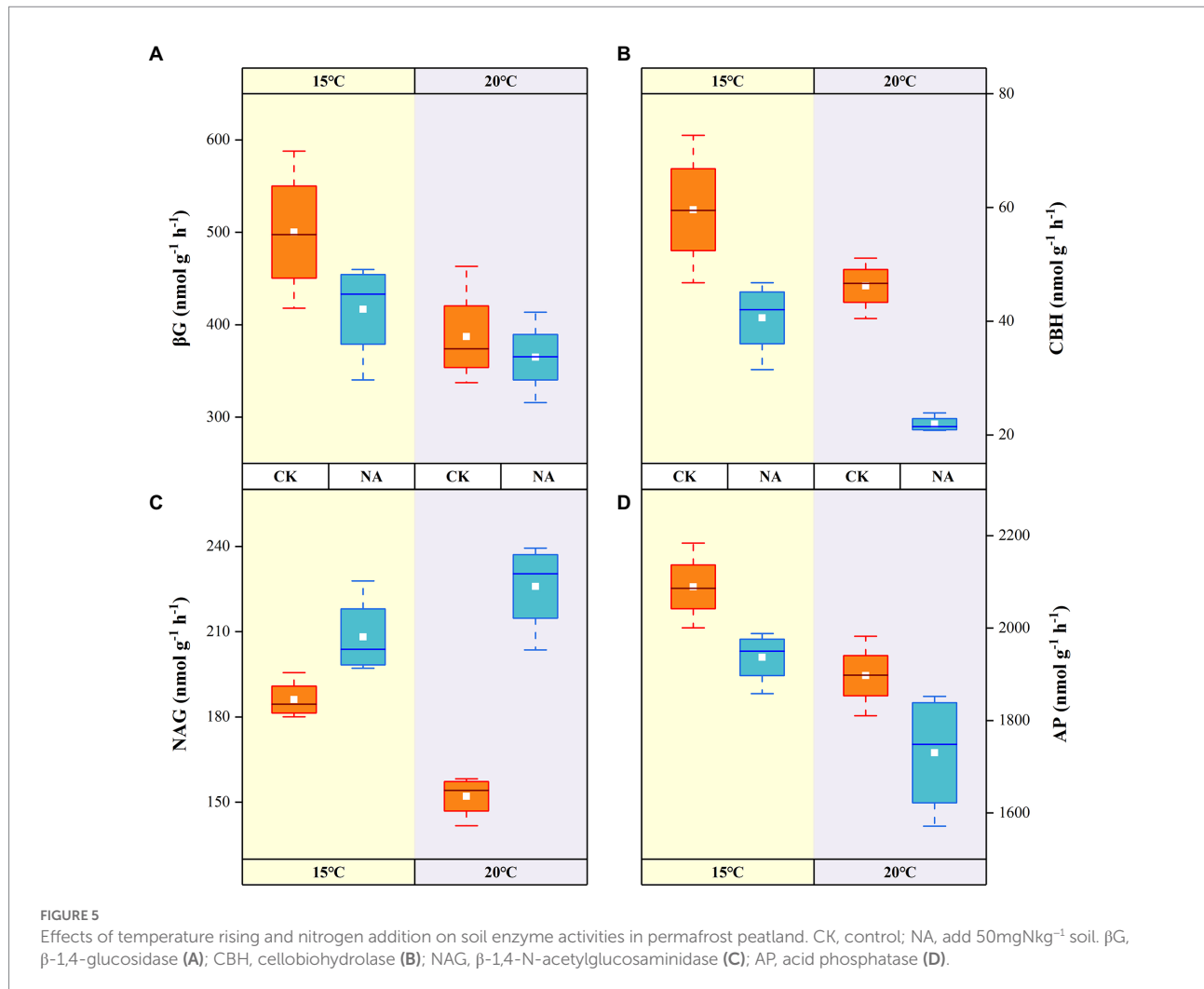
An increase in temperature increased DOC contents in the permafrost peatlands from 531.05 to 628.25 mg kg<sup>-1</sup> in the control treatment (Figure 6A). However, the increase in temperature did not result in a significant change in soil DOC contents under the N addition treatment. N addition had a significantly negative impact on DON contents at 15°C, with DON decreasing from 169.80 to 116.80 mg kg<sup>-1</sup>, whereas soil DON was not significantly affected at 20°C (Figure 6B). NH<sub>4</sub><sup>+</sup>-N in soil ranged from 35.70 to 62.25 mg kg<sup>-1</sup>. Both a rise in temperature and the addition of N resulted in increased contents of soil NH<sub>4</sub><sup>+</sup>-N (Figure 6C). The contents of soil NO<sub>3</sub><sup>-</sup>-N under N addition (19.73 mg kg<sup>-1</sup>) were significantly higher than that in the control (14.96 mg kg<sup>-1</sup>) at 20°C (Figure 6D).

The increase in temperature and N addition had significant interactive impacts on soil DOC and DON contents ( $p < 0.05$ ; Table 1), whereas the effects on NO<sub>3</sub><sup>-</sup>-N and NH<sub>4</sub><sup>+</sup>-N were not significant. The soil emissions of CO<sub>2</sub> and N<sub>2</sub>O showed significant positive correlations with contents of soil NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N, whereas N<sub>2</sub>O emissions were positively correlated with DOC contents ( $p < 0.05$ ; Figure 4).

### Discussion

#### Effect of soil microbial abundances on emissions of soil CO<sub>2</sub> and N<sub>2</sub>O

The rise in temperature and addition of N stimulated emissions of soil CO<sub>2</sub> and N<sub>2</sub>O. Moreover, the rise in temperature

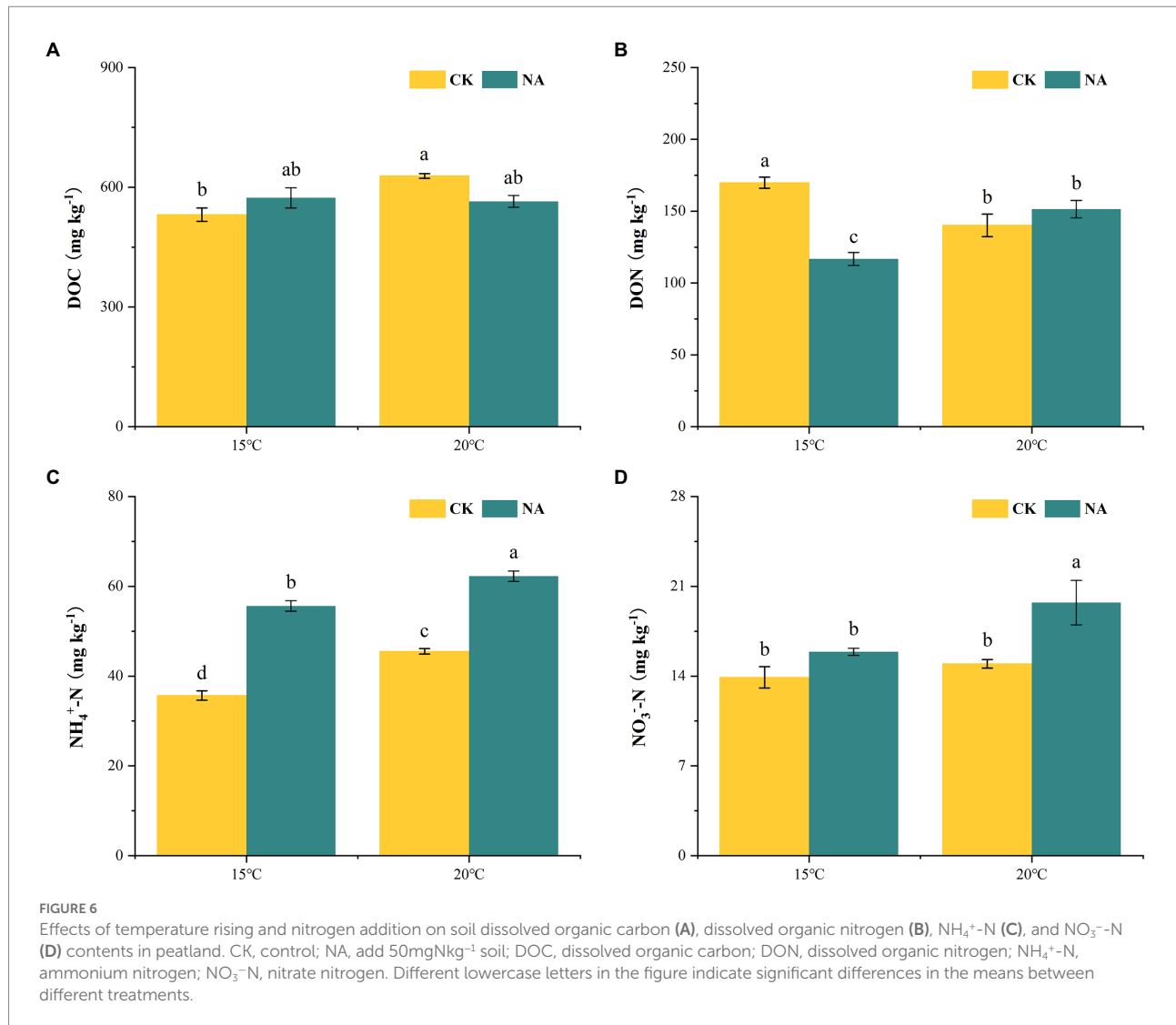


and addition of N interacted within their effect on soil emissions of CO<sub>2</sub> and N<sub>2</sub>O. However, the results of the current study demonstrated a strong negative effect of rise in temperature and the addition of N on the abundances of bacteria. This result indicated that bacteria in permafrost peatlands were adapted to a low temperature and N-limited environment. In line with our results, warming reduced 37% of bacterial abundance and microbial metabolic capacity in the deep organic layer of an Alaska tundra (Wu et al., 2022). Our results showed that the combined effects of temperature rising and N addition significantly increased fungal abundances and there were significantly positive correlations between fungal abundances and the emissions of CO<sub>2</sub> and N<sub>2</sub>O, suggesting that there were differences in sensitivity of different microbial communities to environmental changes and fungi communities played a vital part in the variations of CO<sub>2</sub> and N<sub>2</sub>O emissions at higher temperature and under the addition of N. Consistent with the outcomes of the current study, Xu et al. (2017) determined that fungal tolerance to high temperatures played a significant part in N<sub>2</sub>O emissions.

The results of the present study showed that methanotrophs were more sensitive to a changing temperature and the addition

of N compared to other microbial communities. The higher abundances of *nirK*-type denitrifiers at 20°C compared to at 15°C observed in the present study were consistent with results of previous studies in which the abundances of *nirK* genes were promoted by higher temperatures (Jung et al., 2011; Cui et al., 2016). Declines the abundances of *nirS*-type denitrifiers were observed at 20°C compared to those at 15°C. This result demonstrated that *nirS*-type denitrifiers were better adapted to low temperature conditions. The significant positive correlations between the abundances of *nirK*-type denitrifiers and NH<sub>4</sub><sup>+</sup>-N contents and N<sub>2</sub>O emissions observed in the present study indicated that the increase in emissions of N<sub>2</sub>O could be primarily attributed to the denitrification pathway mediated by *nirK* denitrifiers. Jung et al. (2011) similarly observed an increase in *nirK* genes abundances under both warming and the addition of N. The *nirK* denitrifiers mentioned above are bacterial *nirK*, fungal *nirK* also have clear relevance for N<sub>2</sub>O-producing, future understanding the abundance and distribution of denitrifying fungi may provide new insight into soil N<sub>2</sub>O emissions under various environmental settings (Chen et al., 2016).





## Impacts of soil enzyme activities on emissions of soil $\text{CO}_2$ and $\text{N}_2\text{O}$

Soil enzymes play an important role in the mineralization of soil C and N. Therefore, an improved comprehension of the reaction of soil enzyme activities to increasing temperature and the availability of N is crucial for understanding the mechanisms under which emissions of soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  occur. An increased temperature can alter the nutrient acquisition strategies of microbial communities. This is achieved by changing extracellular enzyme activities through the priming of decomposition of SOM, which leads to increased emissions of  $\text{CO}_2$  from peatlands (AminiTabrizi et al., 2022). NAG participates in N conversion and plays a significant part in the decomposition of nitrogenous substances in soil as it facilitates the degradation of chitin (Liu et al., 2019). Chitin is a major source of soil organic N. The addition of N may affect the decomposition of chitin and peptidoglycan, which, in turn, accelerates the activities of NAG

(Liu et al., 2019). Consistent with the outcomes of the present study, Chen et al. (2018) and Liu et al. (2019) determined that N addition significantly increased the activities of NAG by 5.5% and 56.40–204.78%, respectively. The increase in the activities of NAG can be attributed to soil acidification induced by the addition of N. A decrease in pH was shown to positively affect soil NAG activities (Chen et al., 2018). pH is a key driver for the turnover of organic matter in cold soil, regulatory role of pH needs consideration in the future studies (Leifeld et al., 2013). Although the rise in temperature decreased NAG activities in the control treatment, the increase in NAG activities in the N addition treatment indicated that within the combined effect of an elevated temperature and addition of N, the latter had the dominant effect on soil enzyme activities.

The rise in temperature inhibited the activities of soil  $\beta\text{G}$ , CBH, and AP. This result could be attributed to the decrease in enzyme activities possibly being related to a decrease in substrate (e.g., microbial biomass) availability at elevated temperatures.

Wang J. Y. et al. (2020) determined that enzyme activities reduced with increasing incubation time, suggesting that the responses of enzymes reflected changes in the availability of substrate due to warming. The rate of enzyme production has been shown to decrease as substrate is exhausted. The outcomes of the current study illustrated that the warming stimulation of soil respiration readily depleted hydrolysable substrates during incubation without inputs of C sources. Therefore, decreases in the active pool due to warming can result in microbial C starvation (Metcalfe, 2017; Wang J. Y. et al., 2020). In addition, bacterial conversion of  $\text{NH}_4^+\text{-N}$  to  $\text{NO}_2^-\text{-N}$  in the first step of nitrification can further acidify soils through the release of  $\text{H}^+$  into soil solution. Accelerated acidification, in turn, is an important factor inhibiting soil microbial enzyme activities to acute nutrient amendment (Fatemi et al., 2016). Previous studies have also suggested that a decline in soil enzyme activities was attributable to their more rapid inactivation due to warming can help explain attenuation of the warming impact on mineralization of soil C (Alvarez et al., 2018). Changes in redox conditions driven by temperature can result in abiotic destabilization of Fe-organic matter (phenol) complexes. This is a peatland decomposition pathway that was previously underestimated and can result in increased production of  $\text{CO}_2$  and the accumulation of polyphenol-like compounds that could further inhibit the activities of extracellular enzymes (AminiTabrizi et al., 2022).

## Effect of substrate availability on emissions of soil $\text{CO}_2$ and $\text{N}_2\text{O}$

The emissions of soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  were related to the concentrations of  $\text{NO}_3^-\text{-N}$  and  $\text{NH}_4^+\text{-N}$ . Also, soil emissions of  $\text{N}_2\text{O}$  were related to the concentrations of DOC. Similarly, correlations between the soil  $\text{CO}_2$  release and  $\text{NO}_3^-\text{-N}$  and  $\text{NH}_4^+\text{-N}$  concentrations were revealed by Zhang et al. (2018) in mountain forest and meadow ecosystems. These results indicated that higher substrate availability enhanced the activities of soil microbes, which, in turn, resulted in increased emissions of  $\text{CO}_2$  and  $\text{N}_2\text{O}$ . Soil DOC is composed of low molecular weight organic compounds and drives the growth and activity of microbes by acting as an energy source and a substrate (Wang C. M. et al., 2020). The results of the present study showed an increase in DOC with increasing incubation temperature in the control. An elevated temperature accelerated microbial processes and increased C availability in the control, resulting in higher heterotrophic respiration rates and increased release of  $\text{CO}_2$ . However, soil DOC tended to decrease with the addition of N at a higher incubation temperature, indicating that N addition may limit available C. Warming significantly increased inorganic N ( $\text{NH}_4^+\text{-N}$  and  $\text{NO}_3^-\text{-N}$ ; Table 1) due to higher mineralization and nitrification of TN. The above results are consistent with the earlier study of Yuan et al. (2018), and suggest that warming increases soil N

mineralization. Increase in N mineralization resulted in an increase in soil available N contents with increasing incubation temperature. Higher temperatures have been shown to accelerate the denitrification and nitrification processes (Inclan et al., 2012; Zhang et al., 2016). These processes are major pathways of soil emissions or production of  $\text{N}_2\text{O}$  (Zhang et al., 2018; Li et al., 2019).  $\text{N}_2\text{O}$  emissions due to nitrification accounted for 60–80% of total emissions (Zhang et al., 2020). Therefore, the increased availability of C and N in the soil substrate stimulated  $\text{N}_2\text{O}$  emissions by accelerating N transformation under warming.

In addition to soil temperature, the addition of N had profound influences on the emissions of  $\text{CO}_2$  and  $\text{N}_2\text{O}$ . N addition significantly elevated  $\text{NH}_4^+\text{-N}$  and  $\text{NO}_3^-\text{-N}$ , alleviated microbial N limitation, and promoted soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions, thereby accelerating soil C and N cycling. Menyailoa et al. (2014) similarly found an increase in heterotrophic activity by 20–30% after the addition of N. Increase in the availability of N often accelerates soil denitrification and nitrification processes and results in increased emissions of N-oxide (Davidson et al., 2000; Benanti et al., 2014). Especially, when C are available for microbial activity, N availability will have pronounced impacts on nitrification and denitrification (Lu et al., 2015). Consistent with the result of Guo et al. (2020), the results of the present study showed a positive correlation between DOC and  $\text{N}_2\text{O}$  emissions. This result indicated that both labile C and available N concentrations were the dominant factors influencing the emissions of  $\text{N}_2\text{O}$ . DOC is an important factor regulating denitrification and autotrophic and heterotrophic nitrification (Ferrarini et al., 2017). DOC concentrations influence the emissions of greenhouse gasses by regulating microbial metabolism, whereas soil ammonium and nitrate do not have the same regulatory function (Chen et al., 2020). Increased C availability enhances microbial activity, and, in turn,  $\text{O}_2$  consumption, which may lead to sub-aerobic microsites facilitating  $\text{N}_2\text{O}$  emissions by denitrification and nitrifier denitrification (Ma et al., 2022). Consistent with the outcomes of the current study, Zhu et al. (2016) concluded that the sensitivity of soil respiration to temperature was not influenced by the addition of N, indicating that the availability of C substrate may be more important than that of N substrate.

## Conclusion

This study showed that a rise in temperature and the addition of N promoted soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$  emissions. This result implies that future increases in temperature and availability of N will stimulate C and N cycling in the permafrost peatlands. The abundances of fungi were positively correlated with emissions of soil  $\text{CO}_2$  and  $\text{N}_2\text{O}$ , suggesting that fungal communities may play a significant part in driving the exchange of C and N at the soil-atmosphere interface in permafrost peatlands. The abundances of the *nirK*-type

denitrifiers were positively correlated with DOC and  $\text{NH}_4^+\text{-N}$  contents, and emissions of  $\text{N}_2\text{O}$ , suggesting that the denitrification process mediated by *nirK*-type denitrifiers and available substrate may play a significant part in emissions of  $\text{N}_2\text{O}$ . The activities of soil NAG increased with the addition of N and a rise in temperature, and were positively correlated with soil  $\text{CO}_2$  emissions. This result indicated that the activities of soil NAG are more important than those of other enzymes for regulating  $\text{CO}_2$  emissions. The results of the current study improve understanding of how temperature and N availability regulate soil emissions of greenhouse gasses in permafrost peatlands. However, a laboratory study cannot completely reflect the actual response of greenhouse gasses to global warming, and future research should focus on how plants and their interactions with soil microbes regulate greenhouse gas emissions under field conditions.

## Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

## Author contributions

YS: conceptualization, writing – review and editing, and funding acquisition. XC: methodology, data curation, and writing – review and editing. CS: supervision and funding acquisition. ML, ZL, JG, and XW: writing – review and editing. SG: methodology. All authors contributed to the article and approved the submitted version.

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## Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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## Supplementary material

The Supplementary material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/fmicb.2022.1093487/full#supplementary-material>

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