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# LETTER

# Nitrogen transformations differentially affect nutrient-limited primary production in lakes of varying trophic state<sup>ab</sup>

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# Scientific Significance Statement

The idea that P limitation "evolves" in lakes due to slow, but continuous, inputs of fixed nitrogen (N) over long-time scales has persisted in limnology for decades. This study uses a simple model, along with literature values of nutrient loading, denitrification, and N fixation, to suggest that fixed N may not accumulate in many eutrophic lakes because denitrification removes more N on an annual basis than is replenished through N fixation. This result suggests that eutrophic lakes become increasingly N-limited with time and that increased N inputs will further degrade water quality in those ecosystems.

#### **Abstract**

The concept of lakes "evolving" phosphorus (P) limitation has persisted in limnology despite limited direct evidence. Here, we developed a simple model to broadly characterize nitrogen (N) surpluses and deficits, relative to P, in lakes and compared the magnitude of this imbalance to estimates of N gains and losses through biological N transformations. The model suggested that approximately half of oligotrophic lakes in the U.S.A. had a stoichiometric N deficit, but 72–89% of eutrophic and hypereutrophic lakes, respectively, had a similar N deficit. Although reactive N appeared to accumulate in the most oligotrophic lakes, net denitrification perpetuated the N deficit in more productive lakes. Productive lakes exported reactive N via biological N transformations regardless of their N deficit. The lack of N accumulation through N fixation underscores the need for a modern eutrophication management approach focused on reducing total external nutrient loads, including both N and P.

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The capacity of living organisms to modify the chemistry of their environment has been a central theme in aquatic sciences since Redfield (1958) described the biological influence on ocean stoichiometry. Schindler (1977) expanded this concept using whole-lake experiments to suggest that biological N fixation compensates for nitrogen (N) deficiency relative to phosphorus (P), causing primary production in lakes to "evolve" P limitation. The evolution of P limitation concept has inspired a global view of gradual N accumulation in lakes (Schindler 2012) and oceans (Tyrell 1999), and even influenced the concept of proximate vs. ultimate nutrient limitation of terrestrial primary production (Vitousek et al. 2010). One of the few experiments testing this idea directly is the whole-lake fertilization of Lake 227 in the Experimental Lakes Area of Canada. Lake 227 was fertilized with N and P for 20 yr, and then P only for another 20+ years, to test the effects of these nutrients on primary production. A recent analysis of the Lake 227 experiment suggested

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that short-term fixed N inputs may balance the N pool relative to P on a seasonal basis, but without any new N carryover from one year to the next (Higgins et al. 2017). Thus, rather than evolving P limitation through a decades-long, gradual accumulation of fixed N (Schindler 2012), Lake 227 requires new N via N fixation each year to partially resupply a N pool that has continually declined due to denitrification and/or N burial (Scott and McCarthy 2010; Higgins et al. 2017).

In simple terms, the apparent lack of fixed N carryover in lakes from one growing season to the next can occur if annual denitrification (DNF) rates exceed annual N fixation (Nfix) rates (Paerl et al. 2016). Although the net effects of DNF and Nfix on the N budget in lakes vary on many different temporal and spatial scales, longerterm effects of N transformations on primary production may be considered over annual scales that capture seasonal variability of relevant biogeochemical processes. On an annual scale, DNF (g N m<sup>-2</sup> lake surface area yr<sup>-1</sup>) represents N loss from the total N pool (inorganic N + organic N) to the atmosphere as N<sub>2</sub>. Conversely, annual Nfix (g N m<sup>-2</sup> lake surface area yr<sup>-1</sup>) represents the gain of N into the total N pool from atmospheric N<sub>2</sub>. Thus,

net 
$$N_2$$
 flux = DNF-Nfix (1)

represents the direction and magnitude of N transformations into and out of the total N pool. Lakes with DNF rates exceeding Nfix have a positive net  $N_2$  flux (i.e.,  $N_2$  efflux from total N), and lakes with Nfix rates exceeding DNF have a negative net  $N_2$  flux (i.e.,  $N_2$  influx into total N). Unfortunately, there are very few lakes globally where both DNF and Nfix rates have been quantified simultaneously over a range of temporal scales (but see Grantz et al. 2014), so net  $N_2$  flux is an exceedingly rare measurement. Annual DNF rates are typically greater than Nfix rates in diverse lakes worldwide, with higher DNF rates typically measured in lakes with high total N concentrations (Seitzinger 1988), and higher Nfix rates typically measured in lakes with low total N concentrations and, in some instances, low N:P (Howarth et al. 1988; Bunting et al. 2010). Although N-fixing (diazotrophic) cyanobacteria can dominate seasonally in some lakes and provide new N when external N inputs are low (e.g., Beversdorf et al. 2013; Baker et al. 2018), eutrophic lakes have higher DNF rates than less productive lakes due to high organic matter availability and anoxic conditions in sediments and/or stratified water columns (Finlay et al. 2013). Relating net N<sub>2</sub> flux to the stoichiometric balance of N and P may be useful for understanding the causes of N- or P-limitation, or N + P colimitation (i.e., balanced growth), of primary production in lakes (Paerl et al. 2016).

If we assume that primary production has a stoichiometric optimum (but see Klausmeier et al. 2004), then the imbalance of N relative to P ( $N_I$ ) for primary production can be quantified as either a concentration or flux:

$$N_{\rm I} = N_x - \left(P_x * N : P_{\rm opt}\right) \tag{2}$$

where  $N_x$  is the concentration or flux of total N in or to a lake,  $P_x$  is the concentration or flux of total P in or to a lake, and  $N:P_{opt}$  is the stoichiometric optimum of the N:P ratio for primary production. Regardless of whether the  $N_I$  for primary production is computed as a concentration or flux, negative  $N_I$  values indicate a N deficit relative to P, and positive  $N_I$  values indicate a N surplus relative to P. Considering the relationship between  $N_I$  and net N<sub>2</sub> flux, most lakes are likely either P-limited and losing total N due to a positive net N<sub>2</sub> flux (i.e., N<sub>2</sub> efflux because DNF > Nfix; quadrant I, Fig. 1A) or are N-limited but gaining total N due to a negative net N<sub>2</sub> flux (i.e., N<sub>2</sub> influx because DNF < Nfix; quadrant III, Fig. 1A). Focusing on watershed inputs of total N and P explicitly,  $N_I$  can have units identical to net N<sub>2</sub> flux (g N m<sup>-2</sup> lake surface area yr<sup>-1</sup>). Using this approach, lakes



**Fig. 1.** The relationship between  $N_1$  and net  $N_2$  flux into and out of the total N pool in lakes: (**A**) quadrats represent lakes with a N surplus losing reactive N to atmosphere (I), lakes with a N deficit losing N to atmosphere (II), lakes with a N deficit gaining N from atmosphere (III), and lakes with a N surplus gaining N from atmosphere (IV), (**B**) lakes falling above 1:1 line are experiencing N loss relative to P due to net  $N_2$  flux exceeding  $N_1$  and lakes falling below the 1:1 line are gaining N relative to P due to  $N_1$  exceeding net  $N_2$  flux, and (**C**) four categories of lakes emerge by combining (**A**) and (**B**): (1) lakes with a N surplus that are losing N relative to P, (2) lakes with a N deficit that are losing N relative to P, (3) lakes with a N deficit that are gaining N relative to P, and (4) lakes with a N surplus that are gaining N relative to P.

with  $N_1$  > net N<sub>2</sub> flux are gaining N relative to P annually, while lakes with  $N_1$  < net N<sub>2</sub> flux are losing N relative to P (Fig. 1B).

By combining these two classification approaches (Fig. 1A, B), the two dimensional space can be divided into four categories of lakes (Fig. 1C) based on their  $N_{\rm I}$  and net  $N_2$  flux: (1) lakes with a N surplus from external sources and losing N relative to P; (2) lakes with a N deficit from external sources and losing N relative to P; (3) lakes with a N deficit from external sources and surplus from external sources and gaining N relative to P; and (4) lakes with a N surplus from external sources and gaining N relative to P. The combined approach is important because many lakes may have a negative net  $N_2$  flux, but still not gain N at a rate equivalent to the  $N_{\rm I}$  from external sources (and vice versa).

We used this simple stoichiometric model with continentalscale data from 1964 USA lakes (USEPA 2009, 2016) to explore the potential variability in  $N_{\rm I}$  and net  $N_2$  flux across lakes of differing trophic states. The probabilistic nature of this data set was useful in understanding patterns for lakes in diverse landscapes across the USA (Yuan and Pollard 2018). We developed and calibrated the model using data from a subset of lakes, with a global distribution, for which external N and P input data, as well as DNF and/or Nfix measurements, were available. The overall goal of the study was to explore the role of net  $N_2$ flux in influencing the long-term, gradual accumulation of fixed N in lakes, which theoretically causes the "evolution" of P limitation (Schindler 1977).

#### Methods

The Vollenweider (1968) model was used to derive external P and N input ( $P_s$  and  $N_{s}$ , respectively) estimates from measurements of total P and N concentrations ( $P_c$  and  $N_{c}$ , respectively) for the 33 lakes that had measured values for nutrient concentrations and loads (hereafter, "measured lakes"; Supporting Information Table S1), and for 1964 lakes from the 2007 and 2012 US EPA National Lakes Assessment database that had measured data only for nutrient concentrations (hereafter, "modeled lakes"; Supporting Information Table S2). Briefly, the Vollenweider equation used here required the hydraulic residence time  $(\tau_w)$  and annual water loading rate  $(q_s)$  for each lake. The  $\tau_w$  was unknown for all lakes but was expected to vary between 0.1 and 10 yr. The  $q_s$  was also unknown for all lakes but is a linear transformation of  $\tau_w$  and the mean depth ( $\bar{z}$ ) of a lake. The  $\bar{z}$  for the 33 measured lakes was known (Supporting Information Table S1), but the  $\bar{z}$  for the modeled lakes was unknown. However, the maximum depth (zmax) of the modeled lakes was known (Supporting Information Table S2), and several studies have shown a direct relationship between  $z_{\text{max}}$  and  $\bar{z}$  for lakes at landscape scales. We used the empirical relationship reported by Chow-Fraser (1991) to estimate  $\bar{z}$ from  $z_{\text{max}}$  for each of the modeled lakes.

We conducted a two-step calibration and sensitivity analysis to identify the best choice for  $\tau_w$  for all lakes and to determine the effect of this choice and the prediction of  $\bar{z}$  from  $z_{\text{max}}$  in modeled lakes. That analysis revealed that a  $\tau_w$  of 1 yr minimized the root mean square error (RMSE) of prediction for both  $P_s$  and  $N_s$  in measured lakes and resulted in the most accurate prediction of external nutrient loads. Further, modeled  $\bar{z}$  estimates had a minor effect on the variation of  $P_s$  and  $N_s$  estimates when compared to the variation in nutrient concentrations and ratios in the measured data for both measured and modeled lakes. The supplemental material provides a full description of model development, calibration, and sensitivity analysis in deriving  $P_s$  and  $N_s$  estimates for all lakes.

The  $N_{\rm I}$  for each lake was computed assuming an  $N:P_{\rm opt}$  of 23 (by mass), which is the reported N:P threshold for strictly P-limited lakes in a global analysis of N and P limitation of phytoplankton (Guildford and Hecky 2000). We conservatively used the upper range of colimitation (N:P = 23 by mass). rather than the lower range (N:P = 9), for strict N limitation, because we assumed that even moderate N deficiency experienced by N + P colimited phytoplankton (N:P between 9 and 23) could result in Nfix, which could supplement the N pool and increase the N:P within a year. The annual Nfix for each lake was assumed to be equivalent to the rate of N deficit (negative  $N_{\rm I}$ ) multiplied by an efficiency factor (EF). We assumed that EF, which is the proportion of  $N_{\rm I}$  that could be fixed annually, could vary from 1% to 100% and tested this assumption using measured and modeled Nfix rates for measured lakes (Supporting Information Table S1). Calibration analysis revealed that the RMSE of Nfix predictions decreased with decreasing EF, and we chose a 10% EF as the optimal value for the full analysis (see Supporting Information for details). The annual DNF for each lake was estimated as a proportion of  $N_s$  based on Seitzinger et al. (2006). The efficacy of this model was also tested with measured and predicted values from the measured lakes (Supporting Information Table S1). A complete description of the full modeling approach, as well as a description of the calibration, sensitivity analysis, and model optimization decisions are presented in the Supporting Information.

# Results

 $P_{\rm c}$  and  $N_{\rm c}$  concentrations in measured and modeled lakes each varied by 4–5 orders of magnitude. As shown previously (Downing and McCauley 1992), increasing  $N_{\rm c}$  were not stoichiometrically balanced with increasing  $P_{\rm c}$  (Fig. 2A). The variability in nutrient concentrations was proportional to the variability in  $N_{\rm s}$  and  $P_{\rm s}$  for both the measured and modeled lakes (Table 1). Furthermore, the statistical distributions of  $\bar{z}$ ,  $P_{\rm c}$ ,  $N_{\rm c}$ ,  $P_{\rm s}$ ,  $N_{\rm s}$ ,  $N_{\rm I}$ , Nfix, DNF, and net N<sub>2</sub> flux were similar across the measured and modeled lake data sets (Table 1). Model development and calibration analysis (Supporting Information) demonstrated that assumptions used to derive  $N_{\rm I}$  and net N<sub>2</sub> flux for modeled lakes were valid, and that the resulting model estimates reflect reasonable estimates of these biogeochemical fluxes across the continental USA.



**Fig. 2.** (A) Relationship between total P and total N for measured (red circles) and modeled lakes (gray circles). The reference lines in (A) represent the boundaries for strict N limitation (N:P < 9 by mass) and strict P limitation (N:P > 23 by mass) as described by Guildford and Hecky (2000), (B) the cumulative probability distribution of N<sub>I</sub> for measured and modeled lakes, and (C) the cumulative probability distribution of net N<sub>2</sub> flux for measured and modeled lakes.

The  $N_{\rm I}$  for most lakes ranged from -80 to 40 g N m<sup>-2</sup> yr<sup>-1</sup> (Fig. 2B). The distribution of  $N_{\rm I}$  across all lakes was skewed toward a N deficit (Table 1), but this skewness was more apparent for measured lakes than modeled lakes (Fig. 2B). The net N<sub>2</sub> flux for most lakes ranged from -3 to 8 g N m<sup>-2</sup> yr<sup>-1</sup> (Fig. 2C). The distribution of net N<sub>2</sub> flux across all lakes was strongly skewed toward a positive net N<sub>2</sub> flux (i.e., N<sub>2</sub> efflux because DNF > Nfix; Table 1), with zero measured lakes, and less than 20% of modeled lakes, exhibiting a negative net N<sub>2</sub> flux (i.e., N<sub>2</sub> influx because DNF < Nfix; Fig. 2C).

Sixty percent of the most oligotrophic lakes had some degree of N deficit, but the proportion of oligotrophic lakes (chlorophyll *a* [Chl *a*] < 3.5  $\mu$ g L<sup>-1</sup>) with a N deficit decreased as Chl *a* increased from 1 to 3  $\mu$ g L<sup>-1</sup> (Fig. 3). Thus, oligotrophic lakes approaching the oligotrophic-mesotrophic boundary were evenly distributed between N deficit and N surplus. This pattern reversed as Chl *a* increased from 3 to  $1000 \,\mu g \, L^{-1}$ , and the cumulative proportion of lakes with a N deficit increased from 50% to almost 70% (Fig. 3). According to the model output, and using trophic categories identified by Smith et al. (1999), 51% of oligotrophic lakes (Chl  $a < 3.5 \ \mu g \ L^{-1}$ ), 60% of mesotrophic lakes (Chl  $a = 3.5-9 \ \mu g L^{-1}$ ), 72% of eutrophic lakes (Chl  $a = 9-25 \ \mu g \ L^{-1}$ ), and 89% of hypereutrophic lakes (Chl  $a > 25 \ \mu g \ L^{-1}$ ) were experiencing a N deficit. The average deficit observed in lakes with a negative  $N_{\rm I}$ Ν  $(14.6 \pm 38.8 \text{ g N m}^{-2} \text{ yr}^{-1})$  was approximately six times greater than the average N surplus observed in lakes with a positive  $N_{\rm I}$  $(2.4 \pm 7.1 \text{ g N m}^{-2} \text{ vr}^{-1}).$ 

Surprisingly, the distribution of lakes in the four categories of  $N_{\rm I}$  vs. net N<sub>2</sub> flux were skewed toward category 2 lakes, which had a N deficit from external sources and were losing N relative to P (Fig. 4A). The 1314 category 2 lakes had, on average, the highest total P and Chl *a* concentrations (Fig. 4B). By contrast, the 371 category 4 lakes had a N surplus from external sources, were gaining N relative to P (Fig. 4A), and had the lowest average total P and Chl *a* concentrations (Fig. 4B). Another 279 lakes in category 1 (Fig. 4A) had a N surplus from external sources and were losing N relative to P, with moderate total P and Chl *a* concentrations. No lakes fell into category 3 because this condition was mathematically impossible; Nfix rates would have to exceed the N deficit for a lake, which is impossible because Nfix will not occur when reactive N is available from other sources (Baker et al. 2018).

Although very few lakes worldwide had sufficient data to graph measured  $N_{\rm I}$  vs. net  $N_2$  flux (Fig. 4A), all of these lakes fell into category 2, and none had a negative net  $N_2$  flux. However, some lakes had estimates of  $N_{\rm I}$  and either Nfix or DNF rates (Supporting Information Table S1). DNF rates were strongly proportional to  $N_{\rm s}$  (Fig. 4C) and showed a strong fit with the Seitzinger et al. (2006) model. Similarly, Nfix rates were strongly proportional to the ratio of  $N_{\rm s}$ : $P_{\rm s}$  (Fig. 4D), especially when N:P of nutrient loads was less than the threshold of 23 for strict P limitation (Guildford and Hecky 2000).

#### Discussion

The results of this analysis indicate that biological N transformations, particularly in more productive lakes, may perpetuate N deficient conditions for primary production. This result suggests the potential need for a paradigm shift in limnology. Although some oligotrophic lakes may accumulate N as described by Schindler (1977, 2012), eutrophic lakes may exhibit perpetual N limitation (Scott and Grantz 2013; Grantz et al. 2014; Paerl et al. 2016) caused by biogeochemical transformations that convert biologically

Source/statistic $ar{z}$ (m	rification (UNF			D				
	Total P ) (mg L <sup>-1</sup> )	Total N (mg L <sup>-1</sup> )	$P_{\rm s}$ (g P m <sup>-2</sup> yr <sup>-1</sup> )	$N_{\rm s}$ (g N m <sup>-2</sup> yr <sup>-1</sup> )	$N_{\rm l}$ (g N m <sup>-2</sup> yr <sup>-1</sup> )	Nfix (g N m <sup>-2</sup> yr <sup>-1</sup> )	DNF (g N m <sup>-2</sup> yr <sup>-1</sup> )	Net N <sub>2</sub> flux (g N m <sup>-2</sup> yr <sup>-1</sup> )
Measured data—Support	ing Informatic	on Table S1 ( <i>r</i>	<i>i</i> = 33)					
Minimum 0.	1 0.002	0.05	0.03	0.05	-600	0.00	0.05	0.05
25 <sup>th</sup> percentile 4.	5 0.020	0.50	0.61	7.90	-56	0.02	6.6	3.6
Median 5.	9 0.050	0.73	2.40	23.0	-11	0.63	18	7.3
Mean 2	5 0.105	1.13	4.13	44.0	-51	2.7	18	7.0
75 <sup>th</sup> percentile 9.	8 0.085	1.29	4.30	80.0	-2.0	3.0	24	11
Maximum 29	0.950	7.00	27.0	300	190	18	47	14
Modeled data—EPA/NL <sup>A</sup>	2007, Suppoi	rting Informat	tion Table S2 ( $n = 8$	341)				
Minimum 0.	3 0.001	0.01	< 0.01	0.09	-510	0.00	0.03	-32
25 <sup>th</sup> percentile 1.	2 0.009	0.28	0.06	1.66	-4.7	0.00	0.50	0.31
Median 2.	3 0.025	0.57	0.17	3.65	-0.2	0.02	1.1	0.73
Mean 3.	8 0.126	1.37	0.64	6.67	-8.1	0.89	2.0	1.1
75 <sup>th</sup> percentile 4.	9 0.110	1.24	0.55	7.94	0.8	0.47	2.4	1.7
Maximum 2	2 2.67	26.1	25.2	95.3	31	51	29	24
Modeled data—EPA/NLA	2012, Suppoi	rting Informat	tion Table S2 ( $n = 1$	1123)				
Minimum 0.	4 0.004	0.01	0.01	0.06	-650	0.00	0.02	-57
25 <sup>th</sup> percentile 1.	1 0.021	0.33	0.12	1.82	-7.3	0.01	0.55	0.18
Median 1.	8 0.043	0.65	0.27	3.78	-2.0	0.20	1.1	0.63
Mean 3.	0 0.121	1.18	0.70	6.36	-9.8	1.1	3.9	0.86
75 <sup>th</sup> percentile 3.	5 0.107	1.29	0.61	7.46	-0.1	0.73	2.2	1.5
Maximum 2	2 3.64	54.0	29.5	164	160	65	49	49

Transformation-dependent nitrogen limitation



**Fig. 3.** Relationship between Chl *a* concentrations and the  $N_1$  for the modeled lakes. The cumulative probability of a N deficit (negative  $N_1$ ) or N surplus (positive  $N_1$ ) are also shown as a function of the lake Chl *a* concentration.

reactive N forms to relatively unreactive  $N_2$  gas. Oligotrophic lakes may "evolve" P limitation due to three conditions that increase the N:P ratio: (1) nutrient inputs from pristine landscapes have a relatively high N:P compared to the demands of phytoplankton (Downing and McCauley 1992; Bunting et al. 2010); (2) Nfix adds reactive N seasonally when N deficiencies occur (Scott et al. 2008; Higgins et al. 2017); and (3) DNF does not remove adequate N to induce N limitation due to organic matter limitation and/or lack of anoxic conditions (Finlay et al. 2013).

Our study suggests that these conditions may reverse as lakes become increasingly eutrophic (Paerl et al. 2016). The net N<sub>2</sub> flux may rarely, if ever, support N accumulation relative to P in productive lakes. Insufficient N accumulation relative to P even occurs in lakes with extremely low  $N_{\rm I}$ , where Nfix can represent a significant contribution to the annual N budget. Rather than accumulating fixed N, as many as 80% of USA lakes (categories 1 and 2; Fig. 4A,B) may experience N losses relative to P due to biogeochemical N transformations. Both measured and modeled data support the idea that most lakes experience a positive net  $N_2$  flux (i.e.,  $N_2$  efflux because DNF > Nfix; Fig. 2C), even when external inputs are imbalanced toward N deficiency. Although the large proportion of lakes falling into category 2 lakes was surprising, these results highlight the importance of DNF as a keystone ecosystem process and support the explanation that DNF is responsible for the long-term decline of total N in Lake 227 (Scott and McCarthy 2010; Higgins et al. 2017). Our results



**Fig. 4.** (A) The relationship between  $N_1$  and net  $N_2$  flux for measured and modeled lakes, (B) the distribution of lakes across the categories proposed in Fig. 1C, and the total P and Chl *a* concentrations of the modeled lakes in each category, (C) the relationship between  $N_s$  and DNF for measured lakes, and (D) the relationship between the N:P ratio of external inputs and the Nfix for measured lakes.

also provide a biological mechanism for explaining why P accumulates faster than N in lakes experiencing accelerated eutrophication (Yan et al. 2016).

The most accurate calibration of the model in this study used an EF assumption of 10% Nfix efficiency. The relatively low EF could have been caused by assuming  $N:P_{opt} = 23$ , a global threshold separating strict P limitation and N + P colimitation (Guildford and Hecky 2000). Using that threshold, rather than the threshold for strict N limitation (N:P = 9 by mass), could have over-estimated potential Nfix, which therefore required a lower EF for accurate calibration. The N:P from measured lakes was also more skewed toward N deficiency than modeled lakes (Fig. 2B), indicating that studies focused on Nfix were more likely conducted in lakes with a N deficit.

However, the low efficiency of Nfix observed here can also be explained by ecological constraints other than P availability, such as water column turbulence (Moisander et al. 2002), oxygen supersaturation (Paerl 1990, 2018), light (Scott and Grantz 2013), micronutrient availability (Glass et al. 2010), and/or zooplankton grazing (Hambright et al. 2007). Nfix is an energy intensive process requiring a lengthy genetic cascade, which culminates in heterocyst differentiation and nitrogenase production; indeed, Nfix in cyanobacteria is initiated only when cells detect severe N deficiency (Thiel and Pratte 2014). The biomass yield of phytoplankton relying heavily on Nfix is reduced compared to those growing on combined N sources (Baker et al. 2018). These constraints can also explain the relatively low Nfix efficiency observed in this study and magnify the reality that Nfix may only rarely compensate for N deficiency on annual scales (Scott et al. 2008; Higgins et al. 2017). However, more research is needed to understand the relative importance of each of these constraints and how they may be manifested in lakes across diverse landscapes.

If biological N transformations contribute to perpetual N limitation in productive lakes, then this finding has important implications for our understanding of lake ecology and eutrophication management. Annually recurring N limitation can promote large, often colonial, heterocystous cyanobacteria (Levine and Schindler 1999), which have decreased palatability for zooplankton grazers (DeMott et al. 2001; Ghadouani et al. 2003), resulting in decreased efficiency in the trophic transfer of energy and nutrients. However, many lakes also exhibit annually recurring N limitation but remain dominated by non-N-fixing cyanobacterial taxa (e.g., Paerl et al. 2011; Chaffin et al. 2013; McCarthy et al. 2013), which also have decreased palatability and energy transfer efficiency through the ecosystem (Vanderploeg et al. 2001). Our results support previous analyses (e.g., Paerl et al. 2016) showing that high external N inputs, particularly in lakes with moderate to high P inputs, will support higher biomass production than lakes relying on N fixation (Paerl et al. 2016).

The N:P of some anthropogenic nutrient sources, such as human and animal waste, can be very low relative to demands for primary production (Downing and McCauley 1992). However, the N:P of nonpoint source loading from row-crop agricultural watersheds can be quite high relative to demands for primary production (e.g., Stow et al. 2015), and both N and P reductions are needed to control the harmful effects of accelerated eutrophication in these ecosystems (e.g., Paerl et al. 2016). Our study focused on fluxes based on total N and P concentrations, without considering differences in the reactivity and ecological importance of different chemical forms of N and P, which have important consequences for primary production, phytoplankton community structure, and cyanobacteria toxicity (McCarthy et al. 2009; Donald et al. 2011; Davis et al. 2015; Glibert and Burford 2017; Schoffelen et al. 2018). More research is needed to understand the dynamic relationship between dominant chemical forms, internal recycling mechanisms, and nutrient stoichiometry in lakes. Because P does not have a major gaseous component, legacy P accumulated in lakes may fuel elevated production even after excess P inputs are reduced (Song and Burgin 2017). Likewise, legacy N accumulated in lakes and watersheds may also inhibit recovery from eutrophication in aquatic systems (Duff et al. 2009; Van Meter et al. 2018). Our study shows that reducing N inputs, in addition to P inputs, will likely yield a faster decrease in primary production and phytoplankton biomass because DNF counteracts the accumulation of legacy N. Denitrification efficiency, however, decreases at high external N loading rates (e.g., Mulholland et al. 2008; Gardner and McCarthy 2009), further illustrating the need to reduce external N loads to aquatic systems.

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