1	Dynamic Management Zones for Irrigation Scheduling
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16	Highlights
17	• We used Sentinel 2 NDVI time-series to delineate dynamic management zones (MZ)
18	• Changes in MZ patterns were consistent with soil moisture spatiotemporal variability
19	• Data variance fragmentation was used for daily evaluation of the dynamic MZ designs
20	• Soil moisture data and model forecasts can be used to schedule MZ irrigation
21	

### 22 Abstract

Irrigation scheduling decision-support tools can improve water use efficiency by matching 23 irrigation recommendations to prevailing soil and crop conditions within a season. Yet, little 24 research is available on how to support real-time precision irrigation that varies within-season in 25 both time and space. We investigate the integration of remotely sensed vegetation index time-26 series, soil moisture sensor measurements, and root zone simulation forecasts for in-season 27 28 delineation of dynamic management zones (MZ) and variable rate irrigation scheduling. In a 5.8-29 ha maize field in northeastern Spain, unsupervised classification of 2018 Sentinel 2 vegetation sensing time-series delineated dynamic MZs. The number and spatial extent of MZs changed 30 31 through the growing season. A network of inexpensive soil moisture sensors was used to interpret spatiotemporal changes of Sentinel 2 data. Water content was a significant contributor to changes 32 in crop vigor across MZs through the growing season. Real-time cluster validity function analysis 33 34 provided in-season evaluation of the MZ design. For example, the total within-MZ daily soil moisture relative variance decreased from 85% (early vegetative stages) to below 25% (late 35 36 reproductive stages). Finally, using the Hydrus-1D model, a workflow for in-season optimization of irrigation scheduling and water delivery management was tested. Data simulations indicated 37 that crop transpiration could be optimized while reducing water applications between 11 and 38 28.5% across the dynamic MZs. The proposed integration of spatiotemporal crop and soil moisture 39 data can be used to support management decisions to effectively control outputs of crop  $\times$ 40 *environment* × *management* interactions. 41

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Keywords: Remote sensing; Spatial variability; Temporal variability; Precision agriculture; Soil
moisture; Hydrus-1D

#### 45 1. INTRODUCTION

Irrigated agriculture is essential to global food production, especially because of projected 46 population growth (Döll, 2002). Irrigation water is commonly applied uniformly over an entire 47 field. Yet, field soil water content is typically non-uniform because of spatial variability in soil 48 hydraulic properties (Hawley, 1983), topography (Burt and Butcher, 1985), and vegetation growth 49 (Le Roux et al., 1995). When field spatial variability is significant (Baveve and Laba, 2014; Thorp, 50 2019), modified water management that accounts for variability may improve the cost-51 52 effectiveness of irrigation (Liang et al., 2016; Martini et al., 2017) by increasing for instance water use efficiency and crop yields and decreasing nutrient leaching. 53

54 Precision agriculture seeks to optimize farming operations via site-specific management plans that vary the application of nutrients and water across a field based on variations in soil and crop 55 conditions (Zhang et al., 2002). Management is prescribed over contiguous areas that have 56 57 homogeneous soil properties and crop conditions. These areas are called management zones (MZ). Different clustering methods, including k-mean, ISODATA, and Gaussian Mixture, are available 58 for delineating MZs based on different data sources (Schepers et al., 2004; Martinez-Casasnovas 59 et al., 2012; Galambošová et al., 2014). Commonly, yield maps, topography, remote sensing data, 60 and soil apparent electrical conductivity are used to delineate MZs (Liu et al., 2018; Scudiero et 61 al., 2018; Ohana-levi et al., 2019). Remote sensing crop-canopy data is frequently used in 62 agriculture because it is noninvasive and data can be downloaded without any cost (Fontanet et 63 al., 2018). 64

Several researchers have defined MZs in specific fields with the goal of increasing crop yield
and decreasing water use. Inman et al. (2008) and Schenatto et al. (2015) delineated MZs with
NDVI data and different crop indices. Liu et al. (2018) delineated MZs based on yield and band

vegetation indices maps. Scudiero et al. (2013) argued that spatial information on soil properties 68 known to affect plant growth should guide MZ delineation. They modeled maize yield spatial 69 variability as a function of salinity, texture, carbon content and bulk density, using geospatial 70 apparent soil electrical conductivity and bare soil reflectance measurements as proxies for these 71 soil properties. A similar study was presented by Reyes et al. (2019), in which MZs were defined 72 with NDVI data and complemented with soil properties. Georgi et al. (2018) developed an 73 74 algorithm to delineate MZs automatically based on remote sensing data. However, one of the 75 disadvantages of this algorithm is that it does not work properly on fields with strong timedependent spatial patterns. All the studies cited above consider MZs to be static and assume no 76 77 dynamic pattern during the growing season. In fields where crop spatial patterns change over time, some researchers have advocated for MZ delineation to also be dynamic (Evans et al., 2013; 78 Haghverdi et al., 2015; Cohen et al., 2016; Scudiero et al., 2018). 79

Water content sensors constitute a vital tool for real-time monitoring of water content dynamics in the field. Although sensors monitor water content at a single point, spatial and temporal variations of soil water content and their interactions with crops can be analyzed if several sensors are installed across the field (Biswas and Si, 2011; Biswas, 2014; Yang et al., 2016; Huang et al., 2019). These measurements can provide information about the source of variability between different MZs and aid in their delineation.

In this study, we integrate crop spatial and temporal information from high-resolution remote sensing, soil water sensor data, and numerical model simulations to investigate irrigation scheduling for dynamic management zones. Specifically, we: i.) characterize the spatial and temporal dynamics of crop-soil-water relations, ii.) delineate and evaluate temporally dynamic 90 management zones for variable rate irrigation, and iii.) provide a workflow for in-season
91 optimization of irrigation scheduling and water delivery management.

#### 93 2. MATERIALS AND METHODS

# 94 2.1 Study Site

The research site was a 5.8-ha maize (Zea mays L.) field located in Raïmat, about 170 km west 95 of Barcelona, Spain (Fig.1). The study region has a semi-arid climate. Summer temperatures 96 average 24 °C, with several days above 40 °C. Summer is the dry season, with rainfall of 45 mm. 97 Land use at the study site has changed over the years (Fig. A.1 of Appendix A). Originally, the 98 site was a forest where no tillage occurred. Approximately 30 years ago, the land was converted 99 100 to a vineyard. The topography of the field was modified, with soil being added or removed in various sections, such that the site can now be regarded as having an anthropogenic soil. In 2017, 101 102 one year before this study, grapevines were removed and maize was grown at the site.

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### 104 2.2 Sowing and Irrigation

105 The field was sectioned into four plots that were each sowed with a different maize variety (Fig.1). The varieties were, from west to east: p0937 (DuPont Pioneer, Johnston, IA), d6980 106 (DEKALB Genetics Corporation, Dekalb, IL), p1524 (DuPont Pioneer), and d6780 (DEKALB). 107 All plots were sown on May 3, 2018, with a sowing density of 90000 seeds ha<sup>-1</sup>. Data from the 108 seed companies indicated that the varieties planted on the west and east edges of the field (p0937 109 and d6780, respectively) grow slightly faster than those planted down the center (d6980 and 110 p1524), although all varieties were anticipated to reach full maturity 125-165 days after sowing. 111 Plants started to emerge on May 12, 2018. The site was harvested on September 22, 2018. 112

113 The field was irrigated with a Solid Set sprinkler system (Nelson Irrigation Corporation, Walla 114 Walla, WA) having 15 x 15 m spacing. Water was delivered over 18 irrigation zones at a rate of 115  $6.5 \text{ L} \cdot \text{m}^{-2} \cdot \text{h}^{-1}$ . Total applied water during the season was 679 mm. Irrigation was applied uniformly over the field with scheduling and depths determined using a crop coefficient approach (FAO56).
For most of the site, irrigation ended 115 days after sowing. But, in two 0.3-ha sections located at
the north-east end of the site, irrigation was halted 74 days after sowing due to soil waterlogging.

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# 120 2.3 Soil, Environment, and Crop Measurements

Field data were collected between May and September 2018. Soil moisture, soil and crop parameters, environmental variables, and remote sensing NDVI data were measured. In May 2018, 33 capacitive EC-5 soil moisture sensors (METER Group, Pullman, WA, USA) with an accuracy of  $\pm$  0.03 cm<sup>3</sup>·cm<sup>-3</sup> (Campbell and Devices, 1986) were installed at 11 locations named P1, P2, ..., and P11 (Fig.1). The sensors were installed at 15, 35, and 50 cm depths. Water content data were registered every 30 minutes using an EM5b data logger (METER Group).

At each station, three disturbed soil samples were collected at 0-5, 5-35, and 30-60 cm depth 127 128 for organic matter (OM) and soil texture analyses. The Walkley-Black method was used to measure OM (Nelson and Sommers, 1996), whereas soil particle size distribution was measured 129 according to the gravimetric method (Gee and Bauder, 1986). Particles were categorized into the 130 following size classes: clay (soil particle diameter, D < 0.002 mm), fine silt (0.002 < D < 0.02131 mm), coarse silt (0.02 < D < 0.05 mm) and sand (0.05 < D < 2 mm). Undisturbed soil cores were 132 also collected at the same locations and depths for measuring soil hydraulic properties. The soil 133 water retention curve (SWRC) and unsaturated hydraulic conductivity curve (HCC) were 134 determined using a combination of three laboratory devices: Hyprop, WP4c, and KSat (METER 135 136 Group). The van Genuchten model (van Genuchten, 1980) was fit to the measured curves using the RETC software (van Genuchten MTh, Leij FJ, 1991) to estimate saturated water content ( $\theta_s$ ), 137 residual water content ( $\theta_r$ ), saturated hydraulic conductivity ( $K_s$ ), and the shape parameters  $\alpha$  and 138

*n*. Principal component analysis (PCA) (Abdi and Williams, 2013; Martini et al., 2017) was used
to investigate the relationships between soil texture, OM, bulk density, and hydraulic parameters.
The PCA calculations were done with Statistica 12 (StatSoft Inc. Tulsa, OK, USA).

A weather station consisting of an ECRN-100 rain gauge (METER Group), a cup anemometer 142 (Davis Instruments, Hayward, CA, USA), and PYR pyranometer and VP-4 relative humidity and 143 144 temperature sensors (METER Group) was installed 150 m from the north-east corner of the field. The measured temperature, wind speed, relative humidity, and solar radiation were used to 145 calculate daily reference evapotranspiration  $(ET_0)$  using the Penman-Monteith equation as 146 specified in FAO Irrigation and Drainage Paper No. 56 (Allen et al, 1998; hereafter "FAO56"). 147 The estimated  $ET_0$  was converted into daily water requirements or potential evapotranspiration 148  $(ET_c)$  using the maize crop coefficient  $(k_c)$  from FAO56. Maximum and minimum daily 149 temperature measurements were used to calculate growing degree days (GDD) according to 150 FAO56 and to determinate reference maize growing stages (Ritchie et al., 1997). 151

Remote sensing data from Sentinel 2 were used to determine normalized difference vegetation
index (*NDVI*) (Rouse et al., 1974),

$$NDVI = \frac{(NIR - Red)}{(NIR + Red)} \tag{1}$$

where *NIR* and *Red* are measured reflectance values in the near-infrared and visible red regions, respectively. *NDVI* was used to evaluate spatial variability in the field. Remote sensing data were downloaded with 10-m spatial resolution every 5 days unless there was cloud coverage. The first and last images downloaded were the 15<sup>th</sup> and 135<sup>th</sup> day after sowing. Remote sensing data were processed with the Sentinel application platform (SNAP) software (Zuhlke et al., 2015).

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## 160 2.4 Management Zones Delineation

Sentinel 2 NDVI was used to characterize the spatial variability of crop vigor through the 161 season. A k-means (also known as "fuzzy c-means") unsupervised clustering algorithm (Odeh et 162 al., 2010) was used to classify the NDVI data into temporally dynamic MZs. The Grouping 163 Analysis tool in ArcMap 10.4.1 (ESRI, Rdlands, CA) was used for the MZ delineation. Anytime 164 a new Sentinel 2 NDVI scene was available at the site, a new MZ scheme was delineated. Designs 165 having 2 to 6 MZs were considered. The Calinski-Harabasz criterion (CHC) (Harabasz et al., 166 1974), Eq. (2), was used to evaluate the clusters and MZ delineations and select the optimum 167 number of MZs. The CHC, also known as a pseudo F-statistic, measures the ratio of between-MZ 168 differences and within-MZ similarity. It is formulated as: 169

$$CHC = \frac{BMZSS/(MZn-1)}{WMZSS/(N-MZn)}$$
(2)

where N is the number of pixels, MZn is the number of considered zones, BMZSS is the betweenzones sum of squares, and WMZSS is the within-zone sum of squares. Large *CHC* values indicate high within-MZ homogeneity and between-MZ heterogeneity.

The *NDVI* averages and maximum and minimum values within each MZ were calculated for further comparison between different MZs. MZs were not defined for the beginning of the season (0-20 day after sowing) because plants had not yet germinated or were not big enough to influence *NDVI*, and for the end of the season (beyond 130 days after sowing) because in that period the crop is in a late phenological stage and not irrigated. Differences in soil properties across MZs over time were assessed using a Kruskal-Wallis (Kruskal and Wallis, 1952) rank test (i.e., a nonparametric analysis of variance), calculated with Statistica 12.

Additionally, we considered an alternative static delineation scheme, subdividing the site into
four contiguous fields corresponding to the planted maize varieties. The *CHC* was calculated for

each available *NDVI* scene to compare the variety-based MZ approach to the dynamic *NDVI*based MZ delineation.

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# 185 2.5 Management Zone Available Water

Soil-water status for the MZs was modeled as plant available water (*AW*) (Liang et al., 2016;
Vellidis et al., 2016; Zurweller et al., 2019):

$$AW^{j}(t) = \frac{1}{Z_{\rm T}} \sum_{m} \left( \frac{\theta^{j,m}(t) - \theta_{\rm wp}^{j,m}}{\theta_{\rm fc}^{j,m} - \theta_{\rm wp}^{j,m}} \right) \Delta z^{m}$$
(3)

where  $AW^{j}(t)$  is the profile average available water at monitoring station j and time t, m 188 indexes the measurement depths,  $\Delta z^m$  (cm) is the depth increment associated with the moisture 189 sensor at depth  $m, Z_{\rm T} = \sum \Delta z_m$  (cm) is the total soil profile depth,  $\theta^{j,m}$  (cm<sup>3</sup>·cm<sup>-3</sup>) is soil water 190 content,  $\theta_{wp}^{j,m}$  (cm<sup>3</sup>·cm<sup>-3</sup>) is the wilting point (water content at -1500 kPa), and  $\theta_{fc}^{j,m}$  (cm<sup>3</sup>·cm<sup>-3</sup>) is 191 field capacity (determined using the method of Twarakavi et al., (2009)). The AW for a MZ was 192 defined to be the average AW for all monitoring stations located within the MZ. Note that the MZ 193 design changed over the growing season, so the MZ membership of some stations also changed. 194 In addition to the CHC calculation on the NDVI data, the spatiotemporal variability of AW was 195 also used for in-season evaluation of the dynamic MZ-design. Following Fraisse et al., (2001), we 196 calculated the daily weighted within-MZ AW variance (4), 197

$$S_{MZ_i}^2 = \frac{N_{S_i}N_t}{N_S N_t} \times \frac{1}{N_{S_i}N_t} \sum_{j,k} \left[AW^j(t_k) - \overline{AW_i}\right]^2 \tag{4}$$

where  $S_{MZ_i}^2$  is the daily weighted *AW* variance within management zone *i*; *j* indexes the monitoring stations within management zone *i*; *k* indexes the measurement times during the current day;  $N_{S_i}$  is the number of stations in management zone *i*;  $N_S$ (= 11) is the total number of stations in the field;  $N_t$  (= 48) is the number of measurements per day (every 30 min),  $AW^j$  is defined by (3), and  $\overline{AW_i}$  is the average profile AW across monitoring stations in management zone *i* and measurement times in the current day. The total within-zone variance is equal to the sum of the weighted within-zone variances,  $S^2 = \sum_i S_{MZ_i}^2$ . By comparing  $S^2$  with the total daily field-wide AW variance, it is possible to determine how much was gained in terms of AW uniformity by dividing the field into MZs (Fraisse et al., 2001).

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#### 208 **3. RESULTS**

# 209 3.1 Soil Properties

210 Texture, OM contents and bulk density ( $\rho_b$ ) values measured at each station are reported in Table 1. The soil texture classes (USDA system) of samples taken from the 11 stations were clay 211 loam (42.4 % of samples), loam (42.4%), and silty clay loam (15.2%). Stations on the east side 212 (P1, P6, P7, P11) of the field had, on average, lower sand and higher silt and clay contents than 213 those on the west. Average OM contents ranged between 0.57 and 1.96 %, which is typical for 214 agricultural soils in the region (Romanyà and Rovira, 2011). Fitted and measured parameters for 215 the soil hydraulic properties measured at each station are reported in Table 2. Consistent with the 216 spatial trend in soil texture noted previously, the SWRCs measured on the east side of the study 217 site (stations P1, P6, P7, P11) had lower fitted n values than in the rest of the site. On the wet end 218 of a retention curve, a lower *n* value corresponds to a more gradual transition in water content as 219 pressure head changes. Figure A.2 of Appendix A compares SWRCs observed at stations on the 220 west (P9) and east (P11) sides of the field. 221

The principal component analysis (PCA) indicated that 8 principal components were needed to explain 95% of the variability in the soil dataset. The first three components, PC1 (30.9%), PC2

(18.6%), and PC3 (15.9%), explained around two thirds of the variance in the soil dataset. 224 Particularly, PC1 indicated that clay content clustered (was positively correlated) with  $\theta_{wp}$ ,  $\theta_{fc}$ , and 225  $\alpha$ . The PC1 also indicated that clay content was negatively correlated with sand content,  $\theta_r$ , and n. 226 Further detail about PC1, PC2, and PC3 are reported in Fig. A.3 of Appendix A. 227

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#### 3.2 **Remote Sensing and Dynamic Management Zones Delineation**

The site average, minimum, and maximum NDVI values for each available Sentinel 2 scene 230 are reported in Fig. 2a. The changes in average NDVI generally corresponded to the evolution of 231  $ET_c$  at the site, consistent with reports for maize grown in Mediterranean climates in other studies 232 (Toureiro et al., 2017; Segovia-Cardozo et al., 2019). Figure 2b shows that cumulative input water 233 (irrigation and precipitation) (618 mm) exceeded by 10.2% the site-wide cumulative  $ET_c$  (561 234 mm). At the bottom of Fig. 3, reference growing stages for maize at the site are shown (Ritchie et 235 al., 1997). Varieties at the site took 120 to 130 days to reach maturity. Thus, we considered the 236 reference growing stages to be representative for all maize varieties grown at the site. NDVI and 237  $ET_c$  were low during the early vegetative stages, had maximum values during the late vegetative 238 stage (VT) through the beginning of the reproductive stages (R1-R6), then decreased after R6. The 239 temporal changes of NDVI at the site are comparable to those observed in other studies on maize 240 241 (Viña et al., 2004). In the early vegetative stages (V0 to V5), the NDVI range of each Sentinel 2 scene was narrow. In later vegetative stages and early reproductive stages, the NDVI ranges were 242 much larger, indicating considerable variability in crop status (greenness, health) at the site. 243

Figure 3a shows the spatiotemporal changes of NDVI at the sites. Areas with high and low 244 245 NDVI were observed at the site throughout the growing season. However, the NDVI spatial patterns changed over time, suggesting that homogeneous or static site-specific management may 246

be inadequate to address crop needs over time at this site. Figure 3b shows the dynamic MZ 247 delineation obtained with unsupervised clustering of the NDVI data. Through the growing season, 248 the number of MZs, as well as their spatial distribution, changed. At the beginning of the season, 249 until 50 days after sowing, the CHC indicated that three MZs were optimal for identifying 250 homogeneous zones at the site. The MZ1 covered the north-west side of the site and had the highest 251 NDVI values; the MZ2 had intermediate NDVI and spanned across the south of the site until the 252 45<sup>th</sup> day after sowing and after that over the south-west only. The MZ3 had lower NDV1 values 253 and was initially the north-eastern side of the site, then covered the entire western side of the field 254 at 45 days after sowing. From the 50<sup>th</sup> day after sowing, the CHC indicated that four clusters were 255 best at identifying areas with homogeneous NDVI. MZ1 and MZ2 remained relatively similar to 256 their early season delineations. The MZ4 identified an area of moderately low NDVI at the south-257 eastern portion of the site, whereas MZ3, on the north-eastern side of the site, was characterized 258 259 by the lowest NDVI values. The spatial patterns of the four MZs changed only slightly over time, until the 130<sup>th</sup> day after sowing, when the size of MZ3 increased remarkably while MZ4 decreased. 260 The unsupervised NDVI clustering was compared to dividing the site into four blocks, one for each 261 maize variety. Figure 3c shows the CHC values for NDVI clustering into dynamic MZ and into 262 varietal-based blocks through the growing season. The dynamic MZ-design strategy had larger 263 CHC values for the entire growing season than the variety-block strategy, indicating that the 264 dynamic MZs identified by unsupervised clustering had more homogeneous NDVI than the 265 varietal blocks. 266

Figure 3a shows contrasting *NDVI* values between the eastern and western side of the field, especially visible along the boundary between the d6980 and p1524 varieties. The boundary between the d6980 and p1524 varieties seemed to be a big factor in the determination of the

boundary between eastern (MZ1 and MZ2) and western (MZ3 and MZ4) zones from 55 to 120 270 days after sowing (Fig. 3b). Figure A.1.f of Appendix A shows the p1524 and d6780 varieties 271 doing relatively poorly in July 2018. So, in addition to different soil hydraulic properties on the 272 east side of the field, crop genetics (e.g., pest resistance, germination rate between the varieties) 273 and uneven management (e.g., mechanical sowing, fertilization, soil tillage) could have been 274 contributing factors to the poor performance of the p1524 and d6780 varieties. Changes in MZ 275 276 delineation over time led to some changes in MZ membership for certain soil-water monitoring 277 stations (Table 3). These changes occurred frequently in the early vegetative stages (until 54 days after sowing). No MZ membership change occurred in the late vegetative and reproductive stages. 278 279 The MZs were characterized by contrasting soil properties throughout the season. The MZ had significantly (p<0.05) different PC1 scores throughout the season according to the Kruskal Wallis 280 test: MZ1 and MZ2 were characterized by low PC1 scores, whereas MZ3 and MZ4 were 281 282 characterized by the highest PC1 scores (Fig. A.3 of Appendix A).

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# 3.3 NDVI and Applied Water

Changes in NDVI and AW across MZs are depicted in Fig. 4a (MZ1), 4b (MZ2), 4c (MZ3), 285 and 4d (MZ4). Through the growing season, NDVI in MZ1 and MZ2 was higher than in MZ3 and 286 MZ4. Furthermore, NDVI was slightly higher in MZ1 than in MZ2. Average AW in MZ1 was 287 close to 1 (i.e., water content was near  $\theta_{fc}$ ) throughout the entire growing season. Average AW in 288 MZ2 was greater than 1 at the beginning of the season (until 45 days after sowing) and then very 289 close to 1 through the end of the growing season. Portions of MZ3 and MZ4 had lower NDVI 290 values than MZ1 and MZ2. In these areas, irrigation was likely excessive. AW was considerably 291 higher than 1 for the entire vegetative growth of maize and during the early reproductive stages. 292

Once irrigation was halted in the northeastern corner of the site (i.e., approximately over the area 293 comprised by MZ3) at 74 day after sowing, the AW in MZ3 gradually decreased until the end of 294 the season, while NDVI in MZ3 remained stable. Halting irrigation in the northeastern corner of 295 the site had little-to-no effect on the spatial extent of MZ3 and the other MZs, as shown in Fig. 3c. 296 The analysis of the daily total within-MZ AW variance  $(S^2)$  provided further support for the use 297 of NDVI to identify areas with similar AW conditions at the site. In Fig. 4e, the calculated total 298 MZ variance is normalized by the daily whole-site AW variance. Especially beyond 45 days after 299 sowing (the beginning of the VT growth stage), the normalized within-MZ variance is much less 300 than 1, showing that a large part of the total AW variance was explained by splitting the site into 301 dynamic MZs delineated based on an analysis of NDVI. Fraisse et al. (2001) used yield within-302 zone variance to evaluate soil-derived MZs at the end of the season. Our results suggest that daily 303 AW S<sup>2</sup> could also be used for in-season evaluation of management zone designs. 304

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#### 306 4. DISCUSSION

#### 307 4.1 NDVI and irrigation scheduling simulations

The AW and NDVI time series data show that soil water content was a major factor 308 309 determining NDVI spatiotemporal variability at the site. NDVI is an indicator of maize crop health, and several studies have found positive correlations between NDVI, AW, and crop growth 310 (Scudiero et al., 2014; West et al., 2018). However, those studies were for water scarce conditions. 311 Crop stress and reductions in growth can occur from too much water in the soil profile as well as 312 too little (Feddes et al., 1978). In the current study, where maize was grown under nearly 313 314 waterlogged conditions for most of the growing season (Fig. 4), changes in NDVI and AW between consecutive Sentinel 2 scenes were negatively correlated, with Pearson r equal to -0.64 (MZ1), -315

0.87 (MZ2), -0.79 (MZ3), and -0.83 (MZ4) (all significant at p<0.05). Thus, as has been noted</li>
elsewhere (Shanahan et al., 2008; Long et al., 2015; Quebrajo et al., 2018; Scudiero et al., 2018), *NDVI* data alone should not be used to make irrigation management decisions; *NDVI* (and/or other
plant canopy information) should be integrated with soil information to properly understand plant
processes at a site.

321 With respect to within-season management decisions, one way to make a connection between NDVI-based dynamic management zone delineation and soil conditions would be to use a 322 simulation model to make within-season forecasts of soil and crop conditions for different 323 management options. In the remainder of this paper, we determine a hypothetical optimal irrigation 324 schedule for each growing stage using the simulation/optimization approach developed by 325 326 Fontanet et al. (Vadose Zone Journal, submitted). We first show that a physically based simulation model, Hydrus-1D (Šimůnek et al., 2016), is consistent with NDVI-based zoning by simulating 327 the field experiment and demonstrating agreement between measured AW and simulated available 328 water (SAW), as well as showing a correspondence between simulated transpiration  $(ST_a)$  rates 329 and NDVI. Next, we use the calibrated model to investigate what-if irrigation scenarios, 330 calculating a hypothetical irrigation scheduling table for each dynamic MZ that could have been 331 generated from NDVI within season to guide irrigation. 332

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# 334 4.2 Hydrus-1D Available Water and Transpiration Simulations

The well-known Hydrus-1D model solves the Richards Equation numerically to simulate variably saturated water flow and root water uptake in soils. The model inputs and parameterizations used in our simulations are detailed in Appendix B. Simulations of the experiment for differing monitoring locations all used the same inputs and parameters except for (i.) the soil hydraulic properties, which were measured at each station during the field campaign
(Table 2), and (ii.) the irrigation boundary condition, which differed only for stations P10 and P11
because irrigation was stopped during the experiment.

In Fig. 5, daily observed *AW* for each station is compared with daily-simulated available water (*SAW*). Generally, good agreement between *AW* and *SAW* existed for all stations, although it is acknowledged that the *AW* time courses were relatively non-dynamic. Still, the simulations were done using independently measured hydraulic properties and without any parameter fitting, so the agreement is quite good (modeling details can be found in Appendix B). Missing data towards the end of the season in P7 was due to rodents chewing on the sensor cables.

348 Figure 6 shows the weekly-simulated actual transpiration  $(ST_a)$  at each MZ and the potential transpiration  $(T_p)$  at the site. At MZ1 and MZ2,  $ST_a$  weekly averages were always equal or near 349 the potential transpiration. At MZ3 and MZ4,  $ST_a$  weekly values were remarkably lower than the 350 potential. There was good correspondence between  $ST_a$  and NDVI at each MZ, with a Pearson r 351 of 0.6 (MZ1), 0.51 (MZ2), 0.69 (MZ3), and 0.82 (MZ4). In agreement with the results discussed 352 for NDVI and AW data (section 3.2. Remote Sensing and Dynamic Management Zones 353 Delineation), low  $ST_a$  values at MZ3 and MZ4 were due to waterlogging (root water uptake and 354 transpiration is reduced in the model whenever simulated soil water content exceeds a threshold 355 value; see Appendix B). Stations in MZ3 and MZ4 (see Table 3) had AW and SAW over 1 for 356 most of the growing season (Fig. 4). 357

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# 359 4.3 Irrigation scheduling for within-season decision making

We adopted the method of Fontanet et al. (Vadose Zone Journal, submitted) to investigate optimal irrigation scheduling for dynamic MZs. In this method, irrigation of duration  $\tau$  [T] is

prescribed whenever the soil moisture content decreases below a critical level  $(h^*)$  as indicated by 362 readings from a soil water pressure head sensor(s). The irrigation rate is assumed to be a fixed 363 constant for a given irrigation system. The recommended duration and threshold are determined 364 using a simulation/optimization procedure. Simulations are made using forecasted daily or weekly 365 366 crop water demand (reference ET<sub>0</sub>) and a range of values for the irrigation scheduling parameters, 367  $h^*$  and  $\tau$ . The optimal parameter values are those that maximize seasonal transpiration in the simulations (transpiration being, for many agronomically important crops, proportional to 368 marketable yield). In adapting the simulation/optimization method, we make separate 369 recommendations for each MZ, and update them whenever there is a change in MZ station 370 membership. The recommended values of  $h^*$  and  $\tau$  for a given MZ are the average values 371 determined for monitoring stations within the zone. For simplicity, we use in this example the 372 known daily potential ET<sub>0</sub> for the forecasted model boundary condition (rather than historical data 373 which would be necessary for actual within-season forecasts). Also, as the season progressed, we 374 triggered irrigation based on readings from progressively deeper sensors. In principle, when 375 multiple sensor depths are available, the sensor depth could be treated as an additional optimization 376 377 parameter. Full details on our implementation of the Fontanet et al. (Vadose Zone Journal, 378 submitted) procedure are given in Appendix C.

Although the Fontanet et al. (Vadose Zone Journal, submitted) method prescribes an optimized irrigation schedule, in practice a grower may not be able to irrigate exactly according to a schedule and sensor readings, particularly when there are multiple management zones. Therefore, we also calculated recommended irrigation durations (or, equivalently, irrigation amounts) for soils that have become dryer than the "optimal" irrigation trigger point.

The resulting irrigation scheduling calendar for dynamic-MZ irrigation is presented in Table 384 4. Optimal irrigation strategies for each growth stage are shown in bold. The other table entries 385 show irrigation recommendations for field sections that are dryer than the optimal trigger point. 386 Strategies alternatives in order to allow agriculture to readjust irrigation in case that some parts of 387 the field do not follow the optimal irrigation recommendation. As a general term, and following 388 the tendency of this work, there are two main optimal parameter groups for irrigation scheduling; 389 (i) MZ1 with  $\tau = [1.9, 2.6] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30]$  kPa; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30]$  kPa; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30]$  kPa; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, 2.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -30] \text{ kPa}$ ; MZ2 with  $\tau = [1.9, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.0] \text{ h} \cdot \text{d}^{-1}$  and  $h^* = [-23.3, -3.$ 390 18.3, -30] kPa. Here, the intervals reflect the temporal variations of optimal values associated with 391 the different growing stages. These parameters represent medium frequent and short irrigations. 392 (ii) MZ3 with  $\tau=2 \text{ h}\cdot\text{d}^{-1}$  and  $h^*=[-10, -20]$  kPa; MZ4 with  $\tau=[2.0, 2.3] \text{ h}\cdot\text{d}^{-1}$  and  $h^*=[-10, -16.7]$ 393 kPa. They represent very frequent and short irrigations. The other situation that can be in the site 394 is when the pressure head threshold is smaller than the corresponding optimal value. In this case, 395  $\tau$  increases because water consumed might be supplied until to arrive similar values of simulated 396 actual transpiration as the optimal irrigation scheduling. 397 Table 5 compares seasonal transpiration and irrigation simulated with optimal scheduling 398 399 versus the amounts obtained simulating the field experiment. For MZ1 and MZ2, the optimal 400 schedule recommended 11 to 13 % less water and increased transpiration by 5 to 8 %. For MZ3,

29 % less water was recommended, with an increase in transpiration of 24 %. And for MZ4, a 17
% reduction in irrigation corresponded to a massive 53% increase in transpiration. These results
are consistent with our earlier findings and discussion indicating the field was over-irrigated,
especially in MZ3 and MZ4.

405

406 6. CONCLUSIONS

Irrigation scheduling is complicated by the spatial and temporal variability of a number of 407 variables and parameters. In this work, we investigated a workflow for improved precision 408 irrigation scheduling using data from a maize field where four maize varieties were sown. The 409 workflow is based on dynamic MZ delineation with unsupervised NDVI clustering. We found that 410 MZs based on NDVI clustering were better able to statistically represent field variability than MZs 411 412 based on maize variety. Additionally, the optimal number and spatial configuration of the MZs were found to change over the growing season. The highest number of MZs was 4. Management 413 Zones 1 and 2 (MZ1 and MZ2) corresponded to field sections where NDVI values reflected a 414 typical maize crop performance, whereas MZ3 and MZ4 featured relatively low NDVI values 415 indicative of poor maize growth. 416

Soil water content data were analyzed to show that the variation in crop performance was attributable to soil hydraulic properties, soil available water, and over-irrigation. Further, a relationship existed between NDVI and soil available water. The results indicated that soil available water could potentially also be used for, or incorporated into, in-season evaluation of management zone designs.

Lastly, we proposed a method of combining dynamic management zone delineation with 422 Hydrus 1-D model forecasts for irrigation scheduling. The field experiment was first simulated to 423 confirm the model parameterization and demonstrate its consistency with the obtained NDVI and 424 soil water content data. We then used model simulations to determine an optimal zonation and 425 irrigation calendar for different crop growth stages that could have been generated and updated in 426 real time during the season. Simulations with the optimized irrigation schedule produced an 427 increase in transpiration and a decrease in water use as compared to the field trial (which, again, 428 429 was over-irrigated). The improvement was especially remarkable for MZ3 and MZ4, where

430 irrigation was reduced by 28.5 and 16.6 %, and transpiration increased by 23.9 and 52.6 %,431 respectively.

In summary, we note that although NDVI is useful for dynamically delineating management zones, for irrigation scheduling, it is recommended that NDVI be combined with some additional measure of soil conditions. Low NDVI values may be indicative of poor crop performance, but without other information it is not possible to determine the cause nor recommend a remedial irrigation or management practice.

437

# 438 Appendix A. Supplementary figures

439 Supplementary material related to this article can be found, in the online version, at doi: #440

# 441 Appendix B. Hydurs-1D Simulations

Hydrus-1D (Šimůnek et al., 2008, 2016) was used to simulated soil moisture dynamics and
water balance components at each monitoring station. Each simulation spanned 105 days, from the
18<sup>th</sup> to the 123<sup>rd</sup> day after sowing. The 60 cm soil profile consisted of three layers/materials, as
specified in Table 2. Soil hydraulic properties were specified using the van Genuchten-Mualem
model (van Genuchten, 1980) as follows:

$$\theta(h) = \begin{cases} \theta_r + \frac{\theta_s - \theta_r}{(1 + |\alpha h|^n)^m} & h < 0\\ \theta_s & h \ge 0 \end{cases}$$
(B.1)

447 and

$$K(h) = K_s S_e^{1/2} \left[ 1 - \left( 1 - S_e^{1/m} \right)^m \right]^2, \tag{B.2}$$

448 where  $\theta$  (cm<sup>3</sup>·cm<sup>-3</sup>) is the volumetric water content; *h* is the soil water pressure head (cm);  $\theta_s$ 449 (cm<sup>3</sup>·cm<sup>-3</sup>) is saturated water content;  $\theta_r$  (cm<sup>3</sup>·cm<sup>-3</sup>) is residual water content;  $K_s$  (cm·d<sup>-1</sup>) is

saturated hydraulic conductivity; *n* and  $\alpha$  are shape parameters;  $S_e = \frac{\theta - \theta_r}{\theta_s - \theta_r}$ ; and m = 1 - 1/n.

In Hydrus, root water uptake is simulated using a sink term *S* which has three parts, the potential transpiration rate  $(T_p)$  (cm·d<sup>-1</sup>), the root density distribution  $(\beta)$  (cm<sup>-1</sup>), and the dimensionless water stress function  $(\alpha(h))$ :

$$S(h, z, t) = \alpha(h, z, t)\beta(z, t)T_p(t)$$
(B.3)

454 The actual transpiration rate  $(T_a)$  (cm·d<sup>-1</sup>) is calculated by integrating Eq. (B.3) over the root zone 455 L<sub>R</sub>:

$$T_a = \int_{L_R} S(h, z, t) dz = T_p \int_{L_R} \alpha(h, z, t) \beta(z, t) dz$$
(B.4)

456 Root depth was measured twice a month during the field campaign at one location. This457 information was used to parameterize the Hydrus root growth module.

458 Water stress ( $\alpha(h)$ ) was modeled using the Feddes et al. (1978) function:

$$\alpha(h) = \begin{cases} \frac{h - h_4}{h_3 - h_4} & h_3 > h > h_4 \\ 1 & h_2 \ge h \ge h_3 \\ \frac{h - h_1}{h_2 - h_1} & h_1 > h > h_2 \\ 0 & h \le h_4 \text{ or } h \ge h_1 \end{cases}$$
(B.5)

Parameterized by four critical values of pressure head, Eq. (B.5) defines maximal uptake ( $\alpha = 1$ ) when the soil water pressure head is  $h_2 \ge h \ge h_3$ . Water uptake decreases linearly above or below that range ( $h_3 > h > h_4$  or  $h_1 > h > h_2$ ). And uptake is zero when  $h \le h_4$  or  $h \ge h_1$ . According to the Hydrus-1D database, the parameter values for maize are  $h_1 = -1.5$ ,  $h_2 = -3.0$ ,  $h_3 = -60$ . and  $h_4 = -800$ . kPa, respectively. The value of  $h_3$  was allowed to vary as a function of evaporative demand as modeled by Hyrdurs-1D.

Three observation nodes were inserted in the domain at the same depths as the soil moisture 465 sensors, 15, 35 and 50 cm. Soil moisture values simulated at the observation nodes were used to 466 determine the simulated available water (SAW), using the same procedure as with the field data. 467 The potential evaporation and transpiration rates were calculated by partitioning ET<sub>c</sub> into potential 468 469 evaporation ( $E_p$ ) and transpiration ( $T_p$ ) based on the canopy cover fraction ( $\alpha$ ) according to Raes 470 et al. (2010). An atmospheric boundary condition was imposed at the surface and a free drainage condition was used at the bottom. Simulated actual transpiration  $(ST_a)$  and simulated applied 471 irrigation (SAI) results from each station were extracted.  $ST_a$  and SAI were calculated by averaging 472 stations located with the dynamic MZs. 473

474

#### 475 Appendix C. Irrigation Scheduling

Irrigation scheduling was optimized using the methodology developed by Fontanet et al. 476 (submitted). All soil, environmental and crop inputs are the same as described previously for the 477 478 Hydrus-1D simulations (Appendix B). Possible values for the irrigation scheduling parameters were constrained to be  $h^* \in \{-10, -20, \dots, -100 \text{ kPa}\}$  and  $\tau \in \{1, 2, 3, 4 \text{ h} \cdot d^{-1}\}$ . The irrigation 479 rate was constant (6.5 L·h<sup>-1</sup>·m<sup>-2</sup>). The soil depth used to trigger irrigation ( $Z_{tr}$ ) changed during 480 481 the growing season, becoming deeper as the season progressed. Irrigation parameters have been defined at each station and at different crop growing stages (V0-V5, V6-V10, V11-V15, VT, R1-482 R6). The optimal irrigation at each grow stage and MZ are the average values obtained for the 483 484 stations located in the MZ.

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# 645 Figures:

#### 646

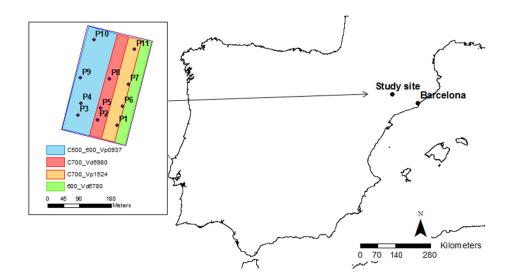




Figure 1. Study site location, soil moisture station locations, and maize variety plantings. The
blue area represents maize variety p0937 (a combination of 500 and 600 series), the red area is
variety d6980 (700 series), the yellow area is p1524 (700 series), and the green area is d6780
(600 series).

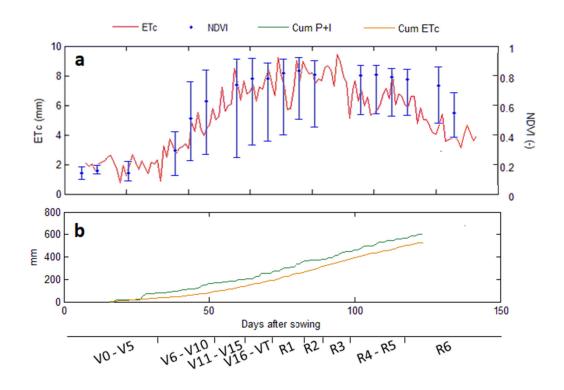




Figure 2. Field average evapotranspiration, NDVI, and cumulative water fluxes as a function
of time and maize growth stage. The bars on the NDVI data indicate field maximum and
minimum values. (V is vegetative stage; R is reproductive stage NDVI is Normalized Difference
Vegetation Index; ETc is daily water requirements; Cum P+I is cumulative Precipitation and
Irrigation; and Cum ETc is cumulative water requirements).

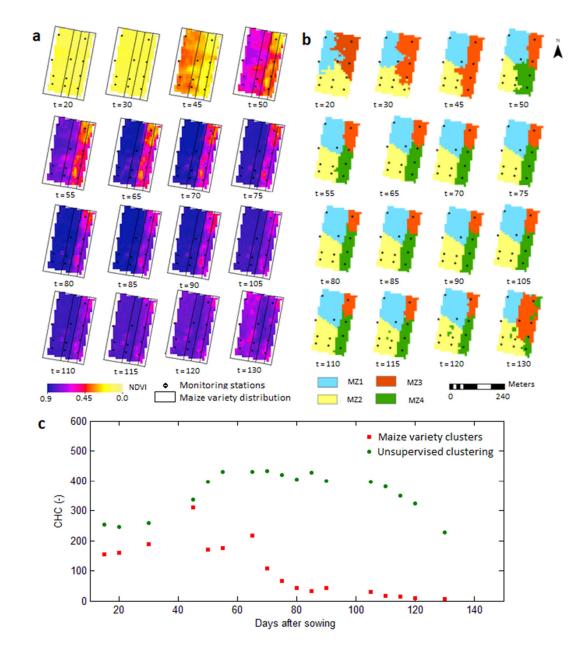


Figure 3. a) Normalized Difference Vegetation Index (NDVI) datasets measured by Sentinel
2 satellite through the growing season; b) dynamic management zone (MZ) delineation. The
letter t indicates days after sowing; and c) Calinski-Harabaz criterion (CHC) for the NDVI
grouped by maize variety (red squares) and with the unsupervised fuzzy-k clustering (green
dots).

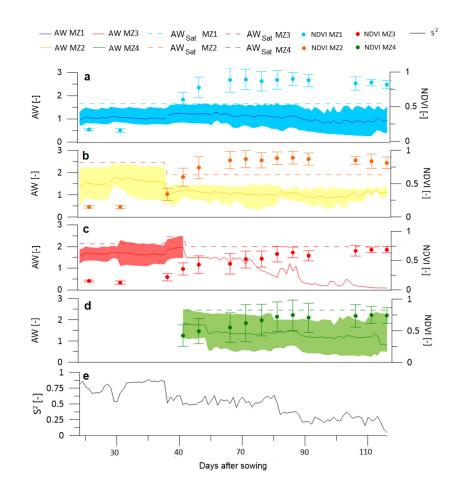


Figure 4. Soil profile available water (AW) and NDVI averages for a) MZ1, b) MZ2, c) MZ3, d) MZ4. Shaded areas represent the maximum in minimum AW at each MZ, while dash lines show available water saturated ( $AW_{sat}$ ) ( $\theta$ ) and field capacity point ( $\theta_{fc}$ ). Error bars represents the maximum and minimum NDVI at each MZ. Note that AW = 1 corresponds to a soil water content equal to field capacity. Panel e) shows the daily total within-MZ weighted variance (S<sup>2</sup>) of AW relative to the daily field-wide AW variance (i.e., S<sup>2</sup>=1).

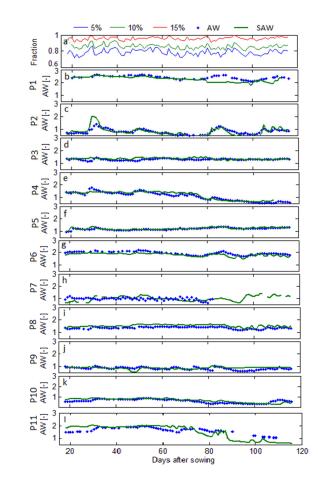


Figure 5. a) Evaluation of profile available water (AW) simulations showings the fraction of
error greater than 5, 10, and 15%. b to l) Comparison between measured available water (AW)
and simulated available water (SAW) at each station (P1 - P11).

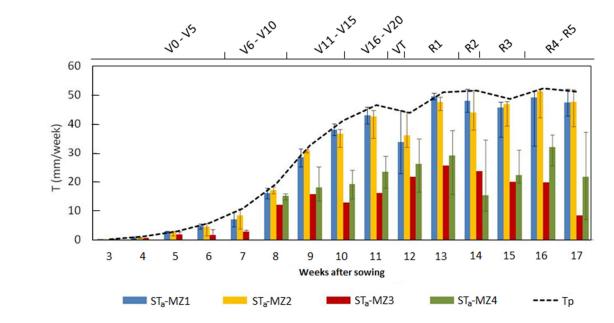


Figure 6. Simulated weekly transpiration at each MZ with the growing stages. Error bars
represent the maximum and minimum and the dash line shows the weekly potential transpiration.

# 685 Tables

Station	Depth	D<0.002 mm	0.002 <d<0.02 mm<="" th=""><th>0.02<d<0.05 mm<="" th=""><th>0.05<d<2 mm<="" th=""><th>OM</th><th><math> ho_{ m b}</math></th></d<2></th></d<0.05></th></d<0.02>	0.02 <d<0.05 mm<="" th=""><th>0.05<d<2 mm<="" th=""><th>OM</th><th><math> ho_{ m b}</math></th></d<2></th></d<0.05>	0.05 <d<2 mm<="" th=""><th>OM</th><th><math> ho_{ m b}</math></th></d<2>	OM	$ ho_{ m b}$
Station	(cm)	Clay (%)	Fine Silt (%)	Coarse Silt (%)	Sand (%)	(%)	$(gr/cm^3)$
P1	0 - 5	36	27.3	13.8	22.9	1.18	1.66
	5 - 35	32	33.6	14.5	19.9	0.71	1.63
	35 - 60	26.5	28.1	9.7	35.7	0.5	1.68
P2	0 - 5	25.9	26.4	14.8	32.9	1.59	1.57
	5 - 35	25.2	26.1	15.1	33.6	1.1	1.58
	35 - 60	24.2	23.4	14.7	37.7	0.98	1.59
Р3	0 - 5	36.5	32.1	14.5	16.9	0.7	1.54
	5 - 35	21.3	27.8	16.7	34.2	0.5	1.65
	35 - 60	24.4	31.8	8.3	35.9	0.65	1.60
P4	0 - 5	28.7	23.6	13.2	34.5	2.71	1.48
	5 - 35	28.5	28.9	11	31.6	1.02	1.59
	35 - 60	28.6	19.8	10.4	41.2	1.14	1.60
P5	0 - 5	22.5	26.3	15.6	35.6	0.57	1.56
	5 - 35	28.9	36.6	20.3	14.2	0.72	1.58
	35 - 60	21.8	28.9	7.3	42.0	0.42	1.56
P6	0 - 5	29.9	26.9	15.1	28.1	2.11	1.64
	5 - 35	29.3	25.7	14.9	30.1	0.85	1.67
	35 - 60	30.2	26	14.8	29.0	0.7	1.69
P7	0 - 5	28.1	36	17.1	18.8	3.14	1.65
	5 - 35	28	27.8	11.9	32.3	1.48	1.72
	35 - 60	27.2	24.3	14.3	34.2	1.27	1.69
P8	0 - 5	25.7	28.7	15.2	30.4	2.22	1.58
	5 - 35	27.7	26.1	14.7	31.5	1.5	1.64
	35 - 60	29.2	27.3	14.7	28.8	1.02	1.78
P9	0 - 5	23.7	26.1	14.8	35.4	2.48	1.53
	5 - 35	23.6	27.8	14.4	34.2	1.06	1.51
	35 - 60	23.5	27.7	14.8	34	0.99	1.51
P10	0 - 5	27.7	25.8	20.3	26.2	1.84	1.61
	5 - 35	28.3	29.5	19.2	26.0	0.72	1.62
	35 - 60	24.6	33.5	9.5	32.4	0.81	1.80
P11	0 - 5	29.4	35.9	14.9	19.8	0.73	1.63
	5 - 35	30.3	34.7	14.9	20.1	0.5	1.65
	35 - 60	26.1	30.5	16.4	27.0	0.5	1.64

**686** Table 1. Soil samples texture, Organic Matter (OM) and bulk density ( $\rho_b$ ) averages at each station.

$\theta_{wp}$ is wiltin								
Station	Depth	$\theta_{s}$	$\theta_{\rm r}$	α	n	Ks	$\theta_{fc}$	$\theta_{wp}$
Station	(cm)	$(cm^{3}cm^{-3})$	$(cm^{3}cm^{-3})$	$(cm^{-1})$	(-)	$(cm \cdot d^{-1})$	$(cm^{3} cm^{-3})$	$(cm^{3} cm^{-3})$
P1	0 - 5	0.424	0.026	0.0169	1.140	2.05	0.345	0.196
	5 - 35	0.407	0.027	0.0150	1.141	2.52	0.351	0.190
	35 - 60	0.364	0.037	0.0115	1.232	1.00	0.350	0.110
P2	0 - 5	0.389	0.061	0.0126	1.364	2.95	0.270	0.103
	5 - 35	0.388	0.060	0.0130	1.358	2.94	0.265	0.104
	35 - 60	0.321	0.047	0.0242	1.354	1.53	0.290	0.124
P3	0 - 5	0.418	0.012	0.0103	1.313	4.42	0.330	0.085
	5 - 35	0.362	0.025	0.0101	1.329	5.63	0.273	0.070
	35 - 60	0.341	0.017	0.0083	1.345	11.47	0.261	0.066
P4	0 - 5	0.439	0.024	0.0658	1.301	5.70	0.340	0.187
	5 - 35	0.400	0.031	0.0143	1.290	4.60	0.300	0.192
	35 - 60	0.395	0.018	0.0424	1.315	4.90	0.315	0.181
P5	0 - 5	0.450	0.062	0.0099	1.497	6.88	0.340	0.070
	5 - 35	0.460	0.067	0.0094	1.402	1.94	0.340	0.080
	35 - 60	-	-	-	-	-	-	-
P6	0 - 5	0.420	0.030	0.0126	1.153	12.00	0.371	0.172
	5 - 35	0.430	0.050	0.0828	1.154	9.40	0.390	0.198
	35 - 60	0.421	0.010	0.0974	1.146	8.10	0.390	0.182
P7	0 - 5	0.375	0.024	0.0105	1.118	1.06	0.300	0.208
	5 - 35	0.349	0.026	0.0380	1.141	3.34	0.300	0.196
	35 - 60	0.361	0.049	0.0391	1.141	4.10	0.280	0.107
P8	0 - 5	0.402	0.040	0.0135	1.375	4.01	0.310	0.123
	5 - 35	0.379	0.030	0.0115	1.356	3.05	0.280	0.090
	35 - 60	0.328	0.020	0.0121	1.287	1.75	0.280	0.080
P9	0 - 5	0.420	0.060	0.0105	1.462	5.79	0.300	0.089
	5 - 35	0.430	0.060	0.0107	1.441	4.70	0.330	0.091
	35 - 60	0.430	0.060	0.0109	1.433	5.52	0.330	0.090
P10	0 - 5	0.389	0.073	0.0115	1.421	3.98	0.301	0.105
	5 - 35	0.387	0.072	0.0112	1.425	4.56	0.300	0.080
	35 - 60	0.320	0.058	0.0181	1.256	1.87	0.290	0.090
P11	0 - 5	0.400	0.012	0.0784	1.121	10.00	0.380	0.188
	5 - 35	0.451	0.018	0.0308	1.141	5.27	0.375	0.188
	35 - 60	0.420	0.014	0.0121	1.112	11.00	0.350	0.250

Table 2. Soil hydraulic parameters from each station, where:  $\theta_s$  is the saturated water content;  $\theta_r$  is the residual water content;  $\alpha$  and n are shape parameters;  $K_s$  is the saturated hydraulic conductivity;  $\theta_{fc}$  is simulated field capacity; and 691  $\theta_{wp}$  is wilting point.

Period (Day after sowing)	MZ1	MZ2	MZ3	MZ4
Period 1 (19-29)	P8, P9, P10	P1, P2, P3, P4, P5, P6	P7, P11	-
Period 2 (30-44)	P8, P9, P10	P1, P2, P3, P4, P5	P6, P7, P11	-
Period 3 (45-49)	P8, P10	P2, P3, P4, P5, P9	P1, P6, P7, P11	-
Period 4 (50-54)	P8, P10	P2, P3, P4, P9	P7, P11	P1, P5, P6
Period 5 (55-115)	P8, P10	P2, P3, P4, P5, P9	P11	P1, P6, P7

Table 3. Periods where one or more stations change MZ membership.

**696** Table 4. Irrigation scheduling calendar based on growing stages and MZs distribution.  $h_{th}$ , is the possible pressure head threshold (the optimal pressure head

697 threshold in bold); *I*, is the irrigation required to maximize transpiration;  $\tau$ , is the irrigation duration;  $Z_{tr}$ , is the trigger soil depth. Optimal irrigation scheduling is 698 represented in bold.

		V0-V5 Trigger Depth = 10 cm		V6-V	/10	V11-	-V15	V	Г	R1-R6		
				Trigger Dep	th = 20cm	00	pth = 20 cm	cm Trigger Depth = 4				
	hth (kPa)	Irrig. Required (mm)	τ (h)	Irrig. Required (mm)	τ (h)	Irrig. Required (mm)	τ (h)	Irrig. Required (mm)	τ (h)	Irrig. Required (mm)	τ (h)	
	0	-	-	-	-	-	-	-	-	-	-	
	-10	-	-	-	-	-	-	-	-	-	-	
	-20	-	-	-	-	-	-	-	-	-	-	
MZ1	-23.3	-	-	12.5	1.9	-	-	-	-	-	-	
	-26.7	13.1	2.0	13.5	2.1	-	-	-	-	-	-	
	-30	14.0	2.2	14.1	2.2	19.0	2.9	15.0	2.3	17.0	2.6	
	-40	15.0	2.3	14.5	2.2	21.0	3.2	17.0	2.6	21.0	3.2	
	-60	16.0	2.5	18.5	2.8	25.0	3.8	23.0	3.5	29.0	4.5	
	-100	17.1	2.6	22.5	3.5	29.0	4.5	33.0	5.1	45.0	6.9	

		Irrig. Required		Irrig. Required		Irrig. Required		Irrig. Required		Irrig. Required	
	hth (kPa)	(mm)	τ (h)								
	0	-	-	-	-	-	-	-	-	-	-
	-10	-	-	-	-	-	-	-	-	-	-
	-18.3	12.4	1.9	-	-	-	-	-	-	-	-
MZ2	-20	12.4	1.9	-	-	-	-	-	-	-	-
	-24	12.9	2.0	15.3	2.4	-	-	-	-	-	-
	-30	13.9	2.1	18.3	2.8	13.0	2.0	13.0	2.0	13.0	2.0
	-40	14.4	2.2	19.3	3.0	18.0	2.8	23.0	3.5	23.0	3.5
	-60	16.9	2.6	24.3	3.7	23.0	3.5	33.0	5.1	33.0	5.1
	-100	19.9	3.1	30.3	4.7	29.0	4.5	45.0	6.9	45.0	6.9

	h <sub>th</sub> (kPa)	Irrig. Required (mm)	τ (h)								
	0	-	-	-	_	-	-	-	_	-	-
	-10	-	-	13.0	2.0	13.0	2.0	13.0	2.0	13.0	2.0
MZ3	-20	13.0	2.0	17.0	2.6	17.0	2.6	21.0	3.2	21.0	3.2
	-30	15.0	2.3	21.0	3.2	21.0	3.2	29.0	4.5	29.0	4.5
	-40	16.0	2.5	23.0	3.5	23.0	3.5	33.0	5.1	33.0	5.1
	-60	18.0	2.8	27.0	4.2	27.0	4.2	41.0	6.3	41.0	6.3
	-100	20.5	3.2	28.0	4.3	28.0	4.3	43.0	6.6	43.0	6.6
	h <sub>th</sub> (kPa)	Irrig. Required (mm)	τ (h)								
	0	-	-	-	-	-	-	-	-	-	-
	-10	-	-	-	-	-	-	13.0	2.0	13.0	2.0
MZ4	-16.7	-	-	-	-	15.0	2.3	15.0	2.3	15.0	2.3
	-20	-	-	-	-	16.0	2.5	17.0	2.6	17.0	2.6
	-30	-	-	-	-	19.0	2.9	25.0	3.8	25.0	3.8
	-40	-	-	-	-	21.0	3.2	29.0	4.5	29.0	4.5
	-60	-	-	-	-	23.0	3.5	33.0	5.1	33.0	5.1
	-100	-	_		-	29.0	4.5	45.0	6.9	45.0	6.9

700Table 5. Comparisons of optimal actual transpiration  $(OpT_a)$ , optimal water applied (OpIA), simulated actual701transpiration  $(ST_a)$ , and simulated water applied (SAI).

1  (- u)	,				
		$OpT_a$	$(OpT_a - ST_a)/ST_a$	OpIA	(OpIA – SAI)/SAI
		(mm)	(%)	(mm)	(%)
	MZ1	405.6	8.0	525.5	-11.0
	MZ2	405.6	4.8	517.8	-12.8
	MZ3	107.5	23.9	217.5	-28.5
	MZ4	271.7	52.6	350.2	-16.6

# 704 Appendix A. Supplementary data

# 



Figure A.1. Historical land use and topography modifications at the study site.

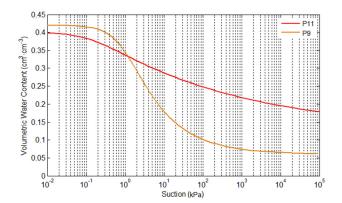
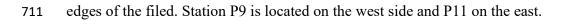


Figure A.2. Soil water retention curves at 10 cm depth from two stations located on opposite



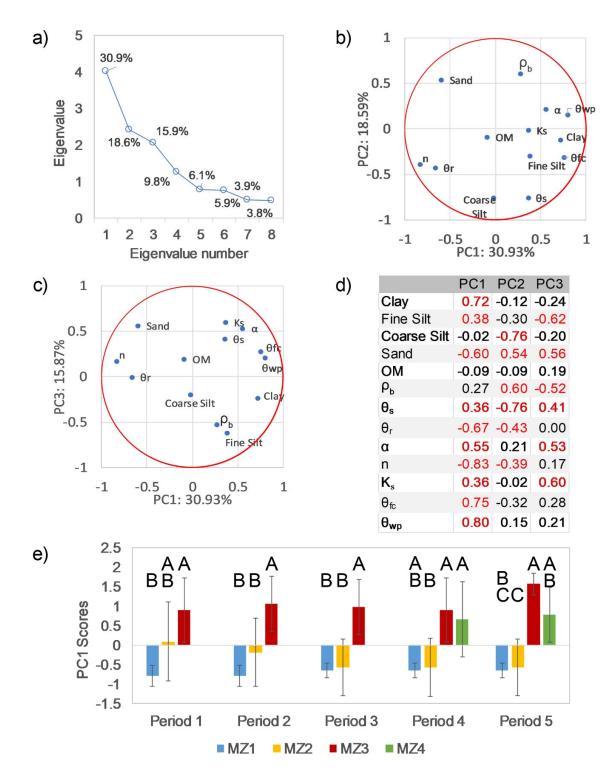


Figure A.3. a) Eigenvalue and percent of variance explained by the first eight components of the
principal component (PC) analysis; b) bi-plot of select soil properties (clay, fine silt, coarse silt,

- sand, and organic matter (OM) content; bulk density ( $\rho_b$ ); water content at saturation ( $\theta_s$ );
- residual water content ( $\theta_r$ ); water content at field capacity ( $\theta_{fc}$ ); water content at wilting point
- 719 ( $\theta_{wp}$ ); saturated hydraulic conductivity ( $K_s$ ); and shape parameters  $\alpha$  and n) for PC1 and PC2; c)
- same for PC1 and PC3; d) Pearson correlation matrix for the first three PCs and selected soil
- 721 properties (significant (p<0.05) correlations highlighted in red); and e) averages (bars) and
- standard deviations (lines) of PC1 for the four management zones (MZs) through the growing
- season. Capital letters indicate significant (p<0.05) differences within MZs according to the
- 724 Kruskal-Wallis test.