THE CONODONT GENUS *TERIDONTUS* (MILLER, 1980) FROM THE EARLY ORDOVICIAN OF MONTAGNE NOIRE, FRANCE

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INTRODUCTION

THE CONODONT genus Teridontus was introduced in 1980 by Miller and was based on the Late Cambrian species Oneotodus nakamurai Nogami, 1967 from the Yencho Member of the Fengshan Fm. of northeast China. Teridontus was later reported from either the Upper Cambrian or Lower Ordovician (Landing et al., 1980; Miller, 1980; Landing and Barnes, 1981; Landing, 1983; An et al., 1983, 1985; Ni et al., 1983; Peng et al., 1983; Nowlan, 1985; Landing et al., 1986; Bagnoli et al., 1987; An, 1987; Buggisch and Repetski, 1987; Pohler and Orchard, 1990; An and Zheng, 1990; Seo and Ethington, 1993; Wang, 1993; Lehnert, 1994; Nicoll, 1994; Seo et al., 1994; Ji and Barnes, 1994; Taylor et al., 1996; Lehnert et al., 1997; Jia, 2000; Dubinina, 2000; Pyle and Barnes, 2002; Zeballo et al., 2005) sediments in numerous localities around the world, but a unanimous interpretation of the composition of the Teridontus apparatus organization was far from accepted.

In 2000 we were invited by J. J. Álvaro, E. P. Villas and D. Vizcaïno to join a working group on the stratigraphy of the Early Paleozoic of southern Montagne Noire (France). Our specific objective was the study of conodont faunas emerging from limestone intercalations in the Ordovician successions. Several samples from the Val d'Homs Formation collected in Sallèles-Caberdès, from Munio Formation, collected in Combes de Barroubio (St. Jean de Minervois) and, chiefly, from the St. Chinian Formation collected in La Regagnade Valley (St. Martial), produced an amazingly rich conodont fauna attributed to the P. deltifer Zone and, more precisely, to the lower part of the P. deltifer deltifer Subzone (Álvaro et al., 2005; Serpagli et al., 2007). The conodont assemblage has also yielded an extraordinarily abundant collection of well preserved elements belonging to a new species of Teridontus that is presented here. A more detailed analysis of the stratigraphic context, distribution and significance of the conodont fauna from southern Montagne Noire is the subject for another study and will be presented at a later date.

THE GENUS TERIDONTUS

Two emended diagnoses of the genus *Teridontus* were provided independently in 1994. Ji and Barnes, describing Early Ordovician conodonts from Western Newfoundland, Canada (1994, p. 64) suggested that four element morphotypes (a, b, c, e) constituted the *Teridontus* multielement structure. According to the authors, the a, b, and c elements (equivalent respectively to Sc, Sb and Sa elements) were respectively symmetrical-subrounded, transitional and suberect-symmetrical whereas the e element (equivalent to M) was compressed. No elements in P positions were identified; furthermore, the early stages of the apparatus might have lacked c, e (Ji and Barnes, 1990, 1994), as well as b elements (Ji and Barnes, 1994).

Nicoll (1994, p. 371), based on a Late Cambrian conodont fauna from Queensland, Australia, regarded *Teridontus* as a seximembrate apparatus bearing two P (Pa and Pb) and four S (Sa, Sb, Sc and Sd) element types, but lacking an element in the M position.

Criteria used to discriminate Teridontus elements in the two

emendations are quite different. Ji and Barnes (1994) referred to the curvature of the cusp (suberect, reclined or recurved), cross section of the cusp and base, and presence of grooves or costae as main features to discriminate apparatus elements. According to Nicoll (1994), the length of the base (short in P and long in S elements) and the cross section of the basal margin were the major features used to segregate discrete element morphologies.

Zeballo et al. (2005), describing Tremadocian conodonts from Argentina, very recently provided a further emendation of *Teridontus*. The authors adopted the Ji and Barnes (1994) apparatus design, which better fitted with their material, but they included also a fifth morphotype (*f*) having a proclined cusp, straight anterior margin and triangular base profile. The difficulty in recognizing Nicoll's six morphotype apparatus structures was possibly attributed by the authors to the Late Cambrian age of Nicoll's and Nogami's original material, compared to the Early Ordovician age of their and of Ji and Barnes' material. On this basis, Zeballo et al. (2005) adopted a "sensu lato" designation for *T. nakamurai*, waiting for new data to come from further investigations.

Löfgren (1997a, 1997b, 1998, 1999a, 1999b) and Löfgren et al. (1999) in recent studies provided detailed reconstructions of several coniform apparatuses [e.g., *Semiacontiodus* (Miller, 1969); *Cornuodus* Fähræus, 1966; *Decoriconus peselephantis* (Lindström, 1955); *Variabiliconus* Landing, Barnes and Stevens, 1986] adopting the apparatus design that Nicoll had also accepted for *Teridontus*. In all such apparatuses, and in *Drepanodus* and *Cordylodus* as well, two P (Pa = g; Pb = f) elements are always present. This fundamental sexi- or septimembrate pattern seems in fact common in latest Cambrian and earliest Ordovician condonts, either in coniform or ramiform apparatuses (e.g., *Cordylodus* Nicoll, 1990; *Junudontus*, Nicoll, 1992; *Oepikodus*, Stewart and Nicoll, 2003, Nicoll and Ethington, 2003; *Erraticodon*, Nicoll and Kelman, 2004). On this basis, we see no reason why these elements should apparently be missing in *Teridontus*.

Our collection from the Montagne Noire, with the recovery of about 700 elements that can be assigned to a new species of *Teridontus*, strongly supports Nicoll's apparatus interpretation, as the six discrete morphologic types have also been recognized in our material. No elements with a laterally compressed cusp or bearing costae, carinae or grooves, diagnostic features of some of Ji and Barnes's morphotypes, have been identified in our specimens. Further considerations are discussed below and in the species diagnosis. The element designations used by Nicoll (1994), Löfgren (1997a, 1997b, 1998, 1999a, 1999b) and Löfgren et al. (1999) are regarded to be of the greatest use in recognising and tracing similarities among the earliest conodont lineages (Löfgren, 1999a, p. 73), and this scheme is followed here.

As regards element curvature, in the Late Cambrian material *Teridontus nakamurai* has the "angle between the posterior margin of the cusp and the upper margin of base . . . essentially a right angle in all elements. The upper part of the cusp may be slightly recurved" (Nicoll, 1994, p. 372). A recent re-analysis of the original material by one of us (RSN) confirmed the absence of any curvature variability in all element types. On the contrary,



FIGURE 1-Element variability in Teridontus gallicus n. sp.

Early Ordovician *Teridontus nakamurai* elements may vary from "erect" (c) to "suberect to reclined" (a) or "reclined to re-curved" (b and e elements) (Ji and Barnes, 1994, p. 65). Elements of *Teridontus gallicus* n. sp. reveal a strong curvature variability in each morphotype of the apparatus, with a gradual variation

from proclined to recurved elements via suberect, erect and reclined elements (Fig. 1). Proclined elements had been previously illustrated in *T. nakamurai* by several authors (Fig. 2). Most of them, interestingly, are Early Ordovician in age. According to our criteria, the newly introduced *f* morphotype



FIGURE 2-Silhouettes of some proclined elements figured by former authors under Teridontus nakamurai. 1, Nowlan, 1985, fig. 5 (27) (earliest Ordovician-Canadian). 2, Wang Cheng-yuan (ed.), 1993, pl. 5, fig. 14 (Late Cambrian). 3, Buggisch and Repetski, 1987, pl. 8, fig. 14 (Tremadocian). 4, Landing and Barnes, 1981, pl. 1, fig. 15 (earliest Ordovician). 5, Bagnoli, Barnes and Stevens, 1987, pl. 2, fig. 18 (early Tremadocian). 6, Miller, 1980, pl. 2, fig. 15 (Late Cambrian). 7, Ji and Barnes, 1994, pl. 24, fig. 3 (earliest Ordovician-Canadian). 8, Dubinina, 2000, pl. 6, fig. 4 (Tremadocian). 9, Jia Cheng-Zao, 2000, pl. 9, fig. 22. 10, Pyle and Barnes, 2002, pl. 15, fig. 18 (earliest Tremadocian). 11, An Tai-Xiang and Zheng Zhao-Cheng, 1990, pl. 1, fig. 3.

of Zeballo et al. (2005) may better fit in Pa (or Pb) positions in the apparatus design here proposed.

SYSTEMATIC PALEONTOLOGY

The described and figured specimens are housed at the Paleontological Museum of the Dipartimento del Museo di Paleobiologia e dell'Orto Botanico of the University of Modena and Reggio Emilia (IPUM) under repository numbers IPUM 27884-27929.

Genus TERIDONTUS Miller, 1980

Type species.—Oneotodus nakamurai Nogami, 1967

Emended diagnosis.—Apparatus composed of six coniform elements differentiated into two P and four S elements. No M element has been recognized. The P elements have a relatively short base, but the bases of the S elements are relatively long. Discrimination of individual P and S types is based on the cross-section of the basal margin. The cusp is essentially round in cross-section, whereas the outline of the base is highly variable between element types. The angle between the upper margin of the base and the posterior margin of the cusp may be highly variable, especially in stratigraphically younger reports. All elements lack grooves, keels, or carinae. Fine striations have been noted on the surface of well-preserved specimens.

TERIDONTUS GALLICUS new species

Figures 1, 3–5

Teridontus nakamurai (Nogami). NOWLAN, 1985, p. 116, fig. 5.26-5.32. Oneotodus? nakamurai Nogami. LANDING, BARNES, AND STEVENS ?1986, pl. 1, figs. 1, 2.

Teridontus nakamurai (Nogami). BUGGISCH AND REPETSKI, 1987, p. 159-160, pl. 2, figs. 1, 3-8, 10-17, 19-21; pl. 3, figs. 1-9, 16-20; pl. 8, figs. 13, 14; pl. 9, figs. 10, 14.

Teridontus cf. nakamurai (Nogami). BUGGISCH AND REPETSKI, 1987, pl. 2, figs. 2, 9, 18; pl. 3, figs. 10-15.

- Teridontus nakamurai (Nogami). BAGNOLI, BARNES, AND STEVENS, 1987, pl. 2, figs. 17, 18; LEE, SEO, AND LEE, 1991, figs. 14, 15, 18.
- Parapanderodus sp. KÜPPERS AND POHLER, 1992, fig. 5.2, 5.7.
- Oneotodus variabilis. KÜPPERS AND POHLER, 1992, fig. 5.8–5.10. Teridontus nakamurai (Nogami). WANG, (Ed.) ?1993, p. 211, pl. 5, figs. 13– 18, 20-23; pl. 8, fig. 25; JI AND BARNES ?1994 (a elements only), pl. 24, figs. 1–3; DUBININA, 2000, p. 196, pl. 6, figs. 2, 4, 5, 7, 8.
- Teridontus nakamurai s.l. (Nogami). ZEBALLO, ALBANESI, AND ORTEGA, 2005, p. 61, fig. 3A-E; Albanesi and Acenolaza, 2005, p. 304, 306, figs. 5 L-M.

Diagnosis.—A species of Teridontus characterized by a gradual variation from proclined to recurved elements via suberect, erect, reclined elements in each element type of the apparatus (Fig. 1). The basal cavity can be observed only in few well preserved elements and is restricted to the base with the pattern outlined by Nicoll (1994, text-fig. 3). All elements lack grooves, keels or carinae. Fine striations have been noted on the surface of wellpreserved specimens (Figs. 3.13a, 4.2b).

Discrimination of individual P and S types is based on the cross-section of the basal margin of the element as follows:

- Pa element (Fig. 1A)—Coniform asymmetrical element with short base obliquely compressed. Outline of basal crosssection irregularly oval with main axis strongly deflected to main axis of the cusp of an angle variable between 30 to 80 degrees.
- Pb element (Fig. 1B)—Coniform asymmetrical element with short base laterally compressed. Outline of basal cross-section irregularly oval with main axis slightly deflected to main axis of the cusp of an angle variable between 10 to 30 degrees.
- Sa element (Fig. 1C)-Coniform symmetrical element with



FIGURE 3—*Teridontus gallicus* n. sp. All elements from sample SM 1. *I*–7, Pa elements; *I*, IPUM 27884, lateral view, $\times 180$; *2*, IPUM 27885, lateral view, $\times 140$; *3*, IPUM 27886, anterior view, $\times 130$; *4*, IPUM 27887, lower-posterior view, $\times 150$; *5*, IPUM 27888, anterior view, $\times 120$; *6*, IPUM 27889, posterior view, $\times 150$; *7*, IPUM 27890, lateral view, $\times 110$. *8–15*, Pb elements. *8*, IPUM 27891, posterior view, $\times 170$; *9*, IPUM 27892, lateral view, $\times 140$; *10*, IPUM 27893, posterior-lateral view, $\times 120$; *11*, IPUM 27894, lateral view, $\times 140$; *12*, IPUM 27895, antero-lateral view, $\times 200$; *13*, IPUM 27896, *a*, fine striations, $\times 750$, *b*, postero-lateral view, $\times 130$; *14*, IPUM 28797, anterior view, $\times 130$; *15*, IPUM 28798, postero-lateral view, $\times 120$.



FIGURE 4—*Teridontus gallicus* n. sp. All elements from sample SM 1. *1–8*, Sa elements. *1*, IPUM 27899, posterior view, ×150; 2, IPUM 27900, *a*, anterolateral view, ×120, *b*, detail of fine striations, ×850; *3*, IPUM 27901, postero-lateral view, ×160; *4*, IPUM 27902, lateral view, ×130; *5*, IPUM 27903, lateral view, ×150; *6*, IPUM 27904, lateral view, ×150; *7*, IPUM 27905, anterior view, ×170; *8*, IPUM 27906, lateral view, ×135. *9–15*, Sb elements. *9*, IPUM 27907, anterior view, ×220; *10*, IPUM 27908, lateral view, ×170; *11*, IPUM 27909, posterior view, ×120; *12*, IPUM 27919, lateral view, ×150; *13*, IPUM 27911, lateral view, ×140; *14*, IPUM 27912, lateral view, ×170; *15*, IPUM 27913, postero-lateral view, ×150.



FIGURE 5—*Teridontus gallicus* n. sp. All elements from sample SM 1. *1–8*, Sc elements. *1*, IPUM 27914, posterior view, ×165; 2, IPUM 27915, lateral view, ×125; 3, IPUM 27916, lateral view, ×105; 4, IPUM 27917, lateral view, ×170; 5, IPUM 27918, lateral view, ×135; 6, IPUM 27919, posterior view, ×150; 7, IPUM 27920, posterior view, ×170; 8, IPUM 27921, antero-lateral view, ×130. *9–15*, Sd elements. *9*, IPUM 27922, anterior view, ×160; *10*, IPUM 27923, lateral view, ×165; *11*, IPUM 27924, postero-lateral view, ×155; *12*, IPUM 27925, lateral view, ×145; *13*, IPUM 27926, postero-lateral view, ×120; *14*, IPUM 27927, anterior view, ×150; *15*, IPUM 27928, lateral view, ×130.

long base dorso-ventrally compressed. Subrounded basal cross-section with main axis perpendicular to main axis of the cusp.

- <u>Sb element</u> (Fig. 1D)—Coniform sub-symmetrical element with long base laterally compressed. Ovate basal crosssection with main axis more or less coincident with main axis of the cusp.
- <u>Sc element</u> (Fig. 1E)—Coniform asymmetrical element with long base irregularly laterally compressed. Irregular ovate basal cross-section having main axis slightly deflected to main axis of the cusp of an angle variable between 10 to 35 degrees.
- <u>Sd element</u> (Fig. 1F)—Coniform asymmetrical element with long base slightly obliquely compressed. Irregularly rounded basal cross-section with main axis more or less deflected to main axis of the cusp of an angle variable between 15 to 60 degrees.

Etymology.-From Gallia, the Latin name for France, area of topotypic material.

Type.—Holotype, IPUM 27890, Figure 3.7. Sample SM 1, Tremadocian (lower part of the *P. deltifer deltifer* Subzone).

Type locality and horizon.—La Regagnade Valley (in the vicinity of St. Martial; lenticular limestone, up to 10 cm thick, located at the lowermost part of the Saint-Chinian Formation, *ca.* 25 m above the top sandstones of the underlying La Dentelle Formation (Álvaro et al., 2003); Early Ordovician (Tremadocian, *Paltodus deltifer deltifer* Subzone of the *P. deltifer* Zone and *Shumardia* (*C.*) *pusilla* Zone).

Other material examined.-679 elements.

Discussion.—The large morphologic variation in the Montagne Noire *Teridontus* population is the basis for this proposal of a new species. Such variability has already been noted by Buggisch and Repetski (1987) under *T. nakamurai* with the illustration of 35 specimens.

As already pointed out, the composition of the apparatus of T. gallicus n. sp. suggested by the Montagne Noire material (six elements types and no M element) is consistent with Nicoll's reconstruction as well as with several Late Cambrian and Early Ordovician coniform apparatuses. However, unlike T. nakamurai in which there is very little morphologic variation of the cusp within each element type, in T. gallicus n. sp. there is a great deal of variation in cusp inclination. The recognition of the element types is not always simple, in part owing to elements with imperfect preservation, especially broken basal margins, but also to the existence of intermediate forms in all P and S element morphotypes. Nevertheless, the great number of morphologies that we include in T. gallicus n. sp. suggests also a need for reconsideration the validity of most of the elements included in the genus Teridontus. In a few specimens, particularly inside the Sa and Sc sequences, the distal part of the cusp may assume some terminal twisting, possibly due to metamorphic deformation. The emended diagnoses presented by Ji and Barnes (1994) and Zeballo et al. (2005) have been already discussed in detail above.

The main differences between *T. gallicus* n. sp. and *T. naka-murai* are as follows:

- 1. Nicoll (1994) has shown that in the large population of Late Cambrian *T. nakamurai* of Australia all elements have an essentially right angle between the posterior margin of the cusp and the upper margin of the base, with only the upper part of the cusp that may be slightly recurved. That means that proclined elements were never found among the 1066 specimen studied. Therefore, the "true" *nakamurai* lacks the variability which characterises our new species.
- Possibly much of the Ordovician *Teridontus* material should be assigned to the new species proposed here, characterised by variability of curvature whereas the latest Cambrian, and

possibly the earliest Ordovician, material remains in *T. nak-amurai*.

3. In *T. nakamurai* the distribution of white matter (albid tissue) in the cusp, except the hyaline tip, is very uniform and with an abrupt transverse and planar transition to a hyaline base, as it is well illustrated by the sketches supplied by Nicoll (1994, fig. 3). The preservation of our material makes it difficult to establish the separation between albid and hyaline tissue.

Comments on the synonymy list are given below.

Nowlan (1985, p. 116) described and illustrated several "markedly reclined" and "strongly proclined" specimens in his fauna. Except for two specimens, all of them were recovered from the Ordovician part of the sections (upper part of Cass Fiord and Cape Clay formations). Possibly also some specimens identified as *T*. aff. *T. nakamurai* belong to *T. gallicus*.

Buggisch and Repetski (1987, p. 159) stressed the great morphological variation in their Early Ordovician material from Antarctica, documented with the illustration of 35 specimens. In our opinion, 14 specimens identified by the authors as *T*. cf. *nakamurai* should be assigned to *T. gallicus* n. sp., thus including the full range of morphologic variability diagnostic of the new species.

Bagnoli et al. (1987) figured two specimens from Tremadocian beds of Western Newfoundland one of which (pl. 2, fig. 18) seems to fit well in the Sb sequence of *T. gallicus* n. sp.

Neither of the two figured specimens reported by Landing, Barnes and Stevens (1986, pl. 1, figs. 1, 2) from the Tremadocian of Quebec is suberect as in typical *T. nakamurai*, but one is proclined and one reclined.

All three specimens from the Early Ordovician of Korea and figured by Lee et al. (1991) fit in the variability of our new species.

Some Montagne Noire specimens figured by Küppers and Pohler (1992) as *Parapanderodus* sp. (fig. 5.2, 5.7) or as *Oneotodus variabilis* (fig. 5.8–5.10) may match our concept of *T. gallicus* n. sp.

In spite of the high variability illustrated by the Chinese specimens reported by Wang (ed., 1993), they are considered with doubts in the synonymy list owing to their Late Cambrian age.

Only a few specimens of Ji and Barnes (1994, pl. 24) may belong to *T. gallicus* n. sp. Much of the material studied by Ji and Barnes probably does not belong within *Teridontus*. The specimens with laterally compressed cusps (Ji and Barnes, 1994, pl. 24, figs. 6–9, 16–17 and 22–25) certainly do not belong in the genus, nor does the element with a lateral groove (pl. 24, fig. 4).

Dubinina (2000) reported several specimens from the Early Ordovician of Central Asia which show high variability and might belong to our new species.

Zeballo et al. (2005) figured two proclined elements (f element of fig. 3E: moderately proclined and c element of fig. 3C: strongly proclined) in their fauna.

Albanesi and Acenolaza (2005) did not illustrate any proclined specimen in their fauna to document the diagnostic curvature variability of our new species. Their material comes from the same formation yielding "*T. nakamurai*" specimens sensu Zeballo et al. (2005) and it is highly probable it may belong to *T. gallicus* n. sp.

Occurrence.—T. gallicus n. sp. has a wide distribution, having being recovered from numerous Early Ordovician localities around the world, including Antarctica (Wright et al., 1984; Buggisch and Repetski, 1987), but, as far as we know, it has been reported in Europe only as *Parapanderodus* sp. and *Oneotodus variabilis* by Küppers and Pohler (1992).

The stratigraphic occurrence of the Montagne Noire elements in the *P. deltifer deltifer* Subzone represents the youngest report of *Teridontus*. This genus, in fact, reaches the *P. deltifer pristinus* Subzone in Argentina (Albanesi et al., 2005; Zeballo et al., 2005) and the claimed occurrence in the lowest Boat Harbour Fm. in Newfoundland, equivalent to *P. deltifer* Zone (Ji and Barnes, 1994, p. 65) needs clarification, as it is not confirmed by distribution reported in tables of pages 91, 92 as well as in text-fig. 21.

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REFERENCES

- ALBANESI, G. L., S. B. ESTEBAN, G. ORTEGA, M. A. HÜNICKEN, AND C. R. BARNES. 2005. Bioestratigrafía y ambientes sedimentarios de las Formaciones Volcancito y Bordo Atravesado (Cámbrico Superior-Ordovícico Inferior), Sistema de Famatina, provincia de La Rioja. Asociación Geológica Argentina, Serie D: Publicación Especial, 8:41–64.
- ALBANESI, G. L. AND G. F. ACENOLAZA. 2005. Conodontes de la Formacion Rupasca (Ordovicico Inferior) en el Angostode Chucalezna, Cordillera Oriental de Jujuy: nuevos elementos bioestratigraficos para una localidad clasica del noroeste argentino. Ameghiniana, 42(2):295–310.
- ÁLVARO, J. J., C. GONZÁLEZ-GÓMEZ, AND D. VIZCAÏNO. 2003. Paleogeographic patterns of the Cambrian-Ordovician transition in the southern Montagne Noire (France): preliminary results. Bulletin de la Société géologique de France, 174:23–31.
- ÁLVARO, J. J., A. FERRETTI, C. GONZÁLEZ-GÓMEZ, C. PIERRE, E. SERPAGLI, I. SUBÍAS, M. VECOLI, AND D. VIZCAÏNO. 2005. An updated Furongian stratigraphic framework for south-western Europe. Acta Micropalaeontologica Sinica, 22(Supplement):3–5.
- AN, T.-X. 1987. The Lower Paleozoic conodonts of South China. Beijing University Press, Beijing, 238 p. [In Chinese]
- AN, T.-X., G.-Q. DU, AND Q.-Q. GAO. 1985. Ordovician conodonts from Hubei, China. Geological Publishing House, Beijing: 1–84. [In Chinese]
- AN, T.-X., F. ZHANG, W. XIANG, Y. ZHANG, W. XU, H. ZHANG, D. JIANG, C. YANG, C. L. LIN, Z. CUI, AND X. YANG. 1983. Conodonts of North China and the adjacent regions. Chinese Science Publishing House, 223 p. [In Chinese]
- AN, T.-X. AND Z.-C. ZHENG. 1990. The conodonts of the marginal areas around the Ordos Basin, North China. Beijing (Science Press), 210 p. [In Chinese with English abstract]
- BAGNOLI, G., C. R. BARNES, AND R. K. STEVENS. 1987. Tremadocian conodonts from Broom Point and Green Point, western Newfoundland. Bollettino della Società Paleontologica Italiana, 35:145–158.
- BUGGISCH, W. AND J. E. REPETSKI. 1987. Uppermost Cambrian (?) and Tremadocian conodonts from Handler Ridge, Robertson Bay Terrane, North Victoria Land, Antarctica. Geologisches Jahrbuch B, 66:145–185.
- DUBININA, S. V. 2000. Conodonts and zonal stratigraphy of the Cambrian-Ordovician boundary deposits. Russian Academy of Sciences, Transactions of the Geological Institute, 517:1–239. [In Russian]
- FÅHRÆUS, L. E. 1966. Lower Viruan (Middle Ordovician) conodonts from the Gullhögen Quarry, southern central Sweden. Sveriges Geologiska Undersökning, Avhandlinger och Uppsatser, C 610:1–40.
- JI, Z. AND C. R. BARNES. 1990. Apparatus reconstructions of Lower Ordovician conodonts from the Midcontinent Province. Courier Forschungs Institut Senckenberg, 118:333–351.
- JI, Z. AND C. R. BARNES. 1994. Lower Ordovician conodonts of the St. George Group, Port au Port Peninsula, western Newfoundland, Canada. Palaeontographica Canadiana, 11:1–149.
- JIA, C.-Z. 2000. Paleozoic stratigraphy and conodonts in Xinjiang. Petroleum Industry Press, 340 p.
- KÜPPERS, A. N. AND S. M. L. POHLER. 1992. Discovery of the first Early Ordovician conodonts from the Montagne Noire, southern France, p. 487– 494. In B. D. Webby and J. R. Laurie (eds.), Global Perspectives on Ordovician Geology, Rotterdam.
- LANDING, E. 1983. Highgate Gorge: Upper Cambrian and Lower Ordovician continental slope deposition and biostratigraphy, northwestern Vermont. Journal of Paleontology, 57:1149–1187.
- LANDING, E. AND C. R. BARNES. 1981. Conodonts from the Cape Clay Formation (Lower Ordovician), southern Devon Island, Arctic Archipelago. Canadian Journal of Earth Sciences, 18:1609–1628.
- LANDING, E., C. R. BARNES, AND R. K. STEVENS. 1986. Tempo of earliest Ordovician graptolite faunal succession: conodont-based correlations from the Tremadocian of Quebec. Canadian Journal of Earth Sciences, 23:1928– 1949.
- LANDING, E., R. LUDVIGSEN, AND P. H. VON BITTER. 1980. Upper Cambrian to Lower Ordovician conodont biostratigraphy and biofacies, Rabbitkettle Formation, District of Mackenzie. Royal Ontario Museum, Life Sciences Contributions, 126:1–42.

- LEE, B.-S., K.-S. SEO, AND H.-Y. LEE. 1991. Conodonts from the Lower Dumugol Shale (Lower Ordovician) of the Cheungsan-Sabuk area, Jeongseon-Gun, Kangweon-Do. Journal of the Paleontological Society of Korea, 1:63–75.
- LEHNERT, O. 1994. A Cordylodus proavus fauna from West-Central Argentina (Los Sombreros Fm., sierra del Tontal, San Juan Province). Zentralblatt für Geologie und Paläontologie, H 1/2:245–261.
- LEHNERT, O., J. F. MILLER, AND J. E. REPETSKI. 1997. Paleogeographic significance of *Clavohamulus hintzei* Miller (Conodonta) and other Ibexian conodonts in an Early Paleozoic carbonate platform facies of the Argentine Precordillera. Geological Society of America Bulletin, 109:429–443.
- LINDSTRÖM, M. 1955. Conodonts from the lowermost Ordovician strata of south-central Sweden. Geologiska Föreningens i Stockholm Förhandlingar, 76:517–604.
- LÖFGREN, A. 1997a. Conodont faunas from the upper Tremadoc at Brattefors, south-central Sweden, and reconstruction of the *Paltodus* apparatus. Geologiska Föreningens i Stockholm Förhandlingar, 119:257–266.
- LÖFGREN, A. 1997b. Reinterpretation of the Lower Ordovician conodont apparatus Paroistodus. Palaeontology, 40:913–929.
- LÖFGREN, A. 1998. Apparatus structure of the Ordovician conodont *Decori*conus peselephantis (Lindström, 1955). Paläontologische Zeitschrift, 72: 337–350.
- LÖFGREN, A. 1999a. The Ordovician conodont Semiacontiodus cornuformis (Sergeeva, 1963) and related species in Baltoscandia. Geologica et Palaeontologica, 33:71–91.
- LÖFGREN, A. 1999b. A septimembrate apparatus for the Ordovician conodont genus *Cornuodus* Fahraeus, 1966. Bollettino della Società Paleontologica Italiana, 37:175–186.
- LÖFGREN, A. M., J. E. REPETSKI, AND R. L. ETHINGTON. 1999. Some translapeticus conodont faunal connections in the Tremadocian. Bollettino della Società Paleontologica Italiana, 37:159–173.
- MILLER, J. F. 1969. Conodont fauna of the Notch Peak Limestone (Cambro-Ordovician), House Range, Utah. Journal of Paleontology, 43:413–439.
- MILLER, J. F. 1980. Taxonomic revisions of some Upper Cambrian and Lower Ordovician conodonts with comments on their evolution. University of Kansas Paleontological Contributions, 99:1–39.
- NI, S.-Z., X.-F. WANG, G.-H. XU, T.-M. ZHOU, Q.-L. ZENG, Z.-H. LI, L.-W. XIANG, AND C.-G. LAI. 1983. Cambrian-Ordovician boundary section at Huanguachang, Yichang, Hubei, China. Bulletin of the Yichang Institute of Geology and Mineral Resources, Chinese Academy of Geological Sciences, 6:79–94.
- NICOLL, R. S. 1990. The genus *Cordylodus* and a latest Cambrian-earliest Ordovician conodont biostratigraphy. Journal of Australian Geology and Geophysics, 11:529–558.
- NICOLL, R. S. 1992. Analysis of conodont apparatus organisation and the genus *Junudontus* (Conodonta), a coniform - pectiniform apparatus structure from the Early Ordovician. BMR Journal of Australian Geology and Geophysics, 13:213–228.
- NICOLL, R. S. 1994. Seximembrate apparatus structure of the Late Cambrian coniform conodont *Teridontus nakamurai* from the Chatsworth Limestone, Georgina Basin, Queensland. AGSO Journal of Australian Geology and Geophysics, 15(3):367–379.
- NICOLL, R. S. AND R. L. ETHINGTON. 2003. Lissoepikodus nudus Gen. et sp. nov. and Oepikodus cleftus sp. nov., new septimembrate conodont taxa from the Early Ordovician of Australia and Nevada. In R. Mawson and J. A. Talent (eds.), Contributions to the Second Australian conodont Symposium (AUSCOS II) held in conjunction with Palaeontology DownUnder—2000 in Orange, Australia, 3–7 July 2000. Courier Forschungsinstitut Senckenberg, 245:427–461.
- NICOLL, R. S. AND A. KELMAN. 2004. Arrangement of elements in the Early Ordovician Likmas type ramiform-ramiform conodont *Erraticodon patu* Cooper, 1981: interpretation and implications. Memoirs of the Association of Australasian Palaeontologists, 30:207–220.
- NOGAMI, Y. 1967. Kambrische Conodonten von China, Teil 2: Conodonten aus den hoch oberkambrischen Yencho-Schichten. Memoirs of the College of Science, University of Kyoto, Geology and Mineralogy, Series B, 33(4): 211–218.
- NOWLAN, G. S. 1985. Late Cambrian and Early Ordovician conodonts from the Franklinian Miogeosyncline, Canadian Arctic Islands. Journal of Paleontology, 59:96–122.
- POHLER, S. M. L. AND M. J. ORCHARD. 1990. Ordovician conodont biostratigraphy, western Canadian Cordillera. Geological Survey of Canada, Paper 90-15:1–37.
- PYLE, L. J. AND C. R. BARNES. 2002. Taxonomy, evolution, and biostratigraphy of conodonts from the Kechika Formation, Skoki Formation, and Road River Group (Upper Cambrian to Lower Silurian), Northeastern British Columbia. NRC Research Press, Ottawa, Ontario, Canada, 227 p.
- SEO, K.-S. AND R. L. ETHINGTON. 1993. Conodonts from the Manitou Formation, Colorado, U.S.A. Journal of the Paleontological Society of Korea, 9:77–92.

- SEO, K.-S., H.-Y. LEE, AND R. L. ETHINGTON. 1994. Early Ordovician conodonts from the Dumugol Formation in the Baegunsan Syncline, Eastern Yeongweol and Samcheog areas, Kangweon-Do, Korea. Journal of Paleontology, 68:599–616.
- SERPAGLI, E., A. FERRETTI, D. VIZCAÏNO, AND J. J. ÁLVARO. 2007. A new conodont genus from the early Ordovician of the southern Montagne Noire, France. Palaeontology. 50(6):1447–1457.
 STEWART, I. AND R. S. NICOLL. 2003. Multielement apparatus structure of the Early Ordovician conodont *Oepikodus evae* Lindström from Australia
- STEWART, I. AND R. S. NICOLL. 2003. Multielement apparatus structure of the Early Ordovician conodont *Oepikodus evae* Lindström from Australia and Sweden. *In* R. Mawson and J. A. Talent (eds.), Contributions to the Second Australian conodont Symposium (AUSCOS II) held in conjunction with Palaeontology DownUnder—2000 in Orange, Australia, 3–7 July 2000. Courier Forschungsinstitut Senckenberg, 245:361–387. TAYLOR, J. F., J. E. REPETSKI, AND C. A. ROEBUCK. 1996. Stratigraphic significance of trilobite and conodont faunas from the Cambrian-Ordovician
- TAYLOR, J. F., J. E. REPETSKI, AND C. A. ROEBUCK. 1996. Stratigraphic significance of trilobite and conodont faunas from the Cambrian-Ordovician shelfbreak facies in the Frederick Valley, Maryland. *In* D. K. Brezinski and J. P. Reger (eds.), Studies in Maryland Geology-In Commemoration of the Centennial of the Maryland Geological Survey. Maryland Geological Survey Special Publication 3:141–164.
- WANG, C.-Y. (ED.).1993. Conodonts of the Lower Yangtze Valley-an index to biostratigraphy and organic metamorphic maturity. Science Press, Beijing, 326 p.
 WRIGHT, T. O., R. J. ROSS, AND J. E. REPETSKI. 1984. Newly discovered
- WRIGHT, T. O., R. J. ROSS, AND J. E. REPETSKI. 1984. Newly discovered youngest Cambrian or oldest Ordovician fossils from the Robertson Bay terrane (formerly Precambrian), northern Victoria Land, Antarctica. Geology, 12:301–305.
- ZEBALLO, F. J., G. L. ALBANESI, AND G. ORTEGA. 2005. Conodontes y graptolitos de las formaciones Alfarcito y Rubasca (Tremadociano) en el área de Alfarcito, Tilcara, Cordillera Oriental de Jujuy, Argentina. Parte 2: Paleontología sistemática. Ameghiniana, 42(1):47–66.
- Leontología sistemática. Ameghiniana, 42(1):47–66.
 ZENG, Q.-L., S.-Z. NI, G.-H. XU, T.-M. ZHOU, X.-F. WANG, Z.-H. LI, C.-G. LAI, AND L.-W. XIANG. 1983. Subdivision and correlation on the Ordovician in the eastern Yangtze Gorges, China. Bulletin of the Yichang Institute of Geology and Mineral Resources, Chinese Academy of Geological Sciences, 6:21–56.

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